

We have revised our manuscript according to the suggestions of the Referee's comments. For clarity, the Referee's comments are reproduced in blue, authors' responses are in black and changes in the manuscript are in red color text. Pages and lines of modified/inserted/deleted texts are relative to the previous version of the manuscript.

Report #1

1. The authors responded that heterogeneous uptake of gas phase aldehydes to the particle phase is responsible for aerosol mass increase in the presence of acid seed aerosol in this study through acid-catalyzed reactions. I am still confused about this part. Is there any direct evidence to support that such reactions do occur in this study? If so, what are the resultant SOA products in the particles? Are they measureable with FTIR or ESI-HRMS? I don't see any supporting data for this statement. The decrease of inorganic sulfate concentration under conditions of high SO₂ concentration is understandable, since formation of organosulfate has been observed, which means that part of the inorganic sulfate contents has been incorporated into the organic molecules.

When studying the effect on acidic seed on the growth of isoprene- and α -pinene-based SOA, it was shown that FTIR peaks at 1180, 1050 and 879 cm⁻¹ are indicators of acid-catalyzed heterogeneous reactions since these peaks could not otherwise be observed in non-acidic conditions (Czoschke et al., 2003 and references therein). In our experiment, although the FTIR peak below 950 cm⁻¹ could not be obtained, two bands similar to 1180 and 1050 cm⁻¹ were observed, which supports the presence of acid-catalyzed reactions in our experiment. Their intensity was very weak when initial SO₂ concentrations were lower than 44 ppb. Although we were not able to identify the SOA products of acid-catalyzed reactions in our experiments, the resulting bands were detected with FTIR.

We added the following at page 9, line 2.

When studying the effect on acidic seed on the growth of isoprene- and α -pinene-based SOA, it was shown that FTIR peaks at 1180, 1050 cm⁻¹ are indicators of acid-catalyzed heterogeneous reactions since these peaks could not otherwise be observed in non-acidic conditions (Czoschke et al., 2003 and references therein). In our experiment, two similar bands located at 1195 and 1040 cm⁻¹, were observed (see Figure S4), which supports the presence of acid-catalyzed reactions in our experiment. These observed peaks are prominent in IR spectra from SOA formation in an acidic particle environment (Jang et al., 2002; Czoschke et al., 2003). In our study, the intensities of these peaks were very weak when initial SO₂ concentrations were lower than 44 ppb, indicating that acid-catalyzed reactions are not facilitated at these conditions.

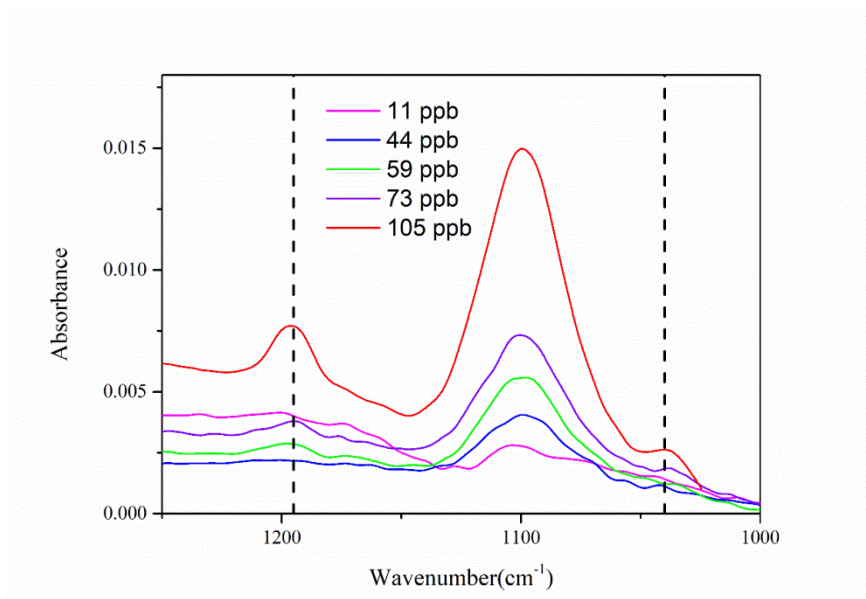


Figure S4: The change of the two peaks at 1195 and 1040 cm^{-1} at different initial SO_2 concentrations.

2. It is still unclear to me about chemical composition of SOA change in the absence versus in the presence of SO_2 . In Figure S5, the authors provided ESI-HR-MS spectra with different initial SO_2 concentrations (236 ppb versus 93.7 ppb), and stated that no obvious composition change under these conditions. However, these are both under high SO_2 conditions. Without the initial input of SO_2 (i.e. initial $\text{SO}_2=0$ ppb), the SOA yield was already significant, but acid-catalyzed reactions and organosulfate formation would not be expected under such conditions. This points to different SOA formation mechanisms that determine the SOA composition and eventually the SOA yield under high SO_2 and zero SO_2 conditions. In this context, it would be more meaningful to present ESI-HR-MS data from 0 ppb, 40 ppb, and 100 ppb to discuss how composition change may affect the SOA yield.

More HR-MS analyses were performed, with 0 ppb and 236 ppb initial SO_2 concentrations (see Figure S6). As can be seen from this Figure, the SOA composition changes with initial SO_2 concentrations.

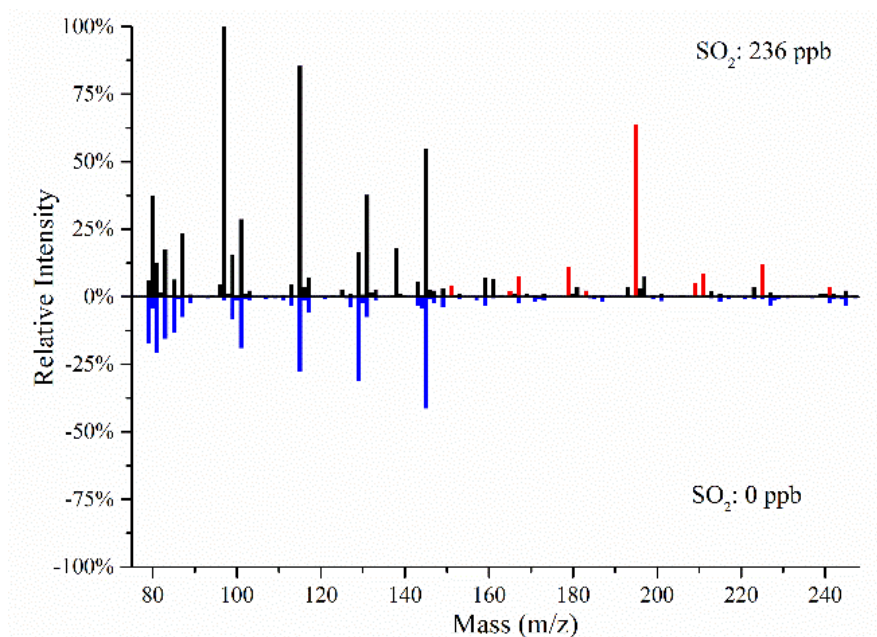


Figure S6: Comparison of SOA ESI-HR-MS spectra with different initial SO₂ concentrations.

We added the following at page 12 line 13.

Comparing HR-MS data when initial concentrations were 0 ppb and 236 ppb reveals that the bands representing organosulfates do not appear at 0 ppb of SO₂. Peaks at m/z larger than 150 were undetectable at initial SO₂ concentration of 0 ppb, while products without sulfur peaked at both concentrations, with the only difference being their intensities. This implies that the process of SOA formation strongly depends on initial SO₂ concentrations.

3. The enhancement of organosulfate yield with increasing initial SO₂ concentrations should be supported with quantitative or semi-quantitative ESI-HR-MS data with authentic or surrogate standards. It is unclear that how the relative intensity of $m/z = 97$ and the peak of organosulfate species could serve as an indicator for increasing organosulfate formation.

Organosulfates could be identified by ESI-HR-MS, but their concentrations should be measured using liquid chromatography (LC). The HR-MS results were obtained through a direct injection as mentioned in page 6 line 4. However, since LC was not combined to HR-MS, quantitative analysis of the SOA chemical composition was not made and surrogate standards were not used. All MS signals could be obtained at the same time, as shown in the figure below.

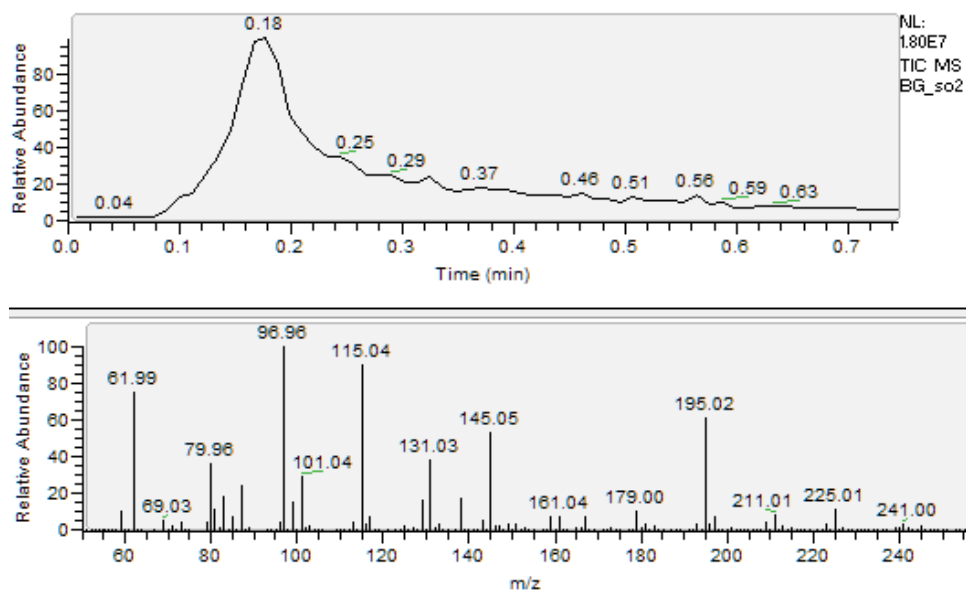


Figure S7: Raw HR-MS data as shown in the Xcalibur 2.2 software. The total ion chromatographic prints of ESI-HR-MS (top) and the mass spectra (bottom).

The HR-MS signals of organosulfates were related to those of sulfates as shown in Figure S6 and S7, which means the concentrations ratios of organosulfates and sulfates were consistent. Here, the amount of sulfates formed increased with increasing initial SO₂ concentration. Since the relative intensities of sulfates and organosulfates are similar, it follows that organosulfates concentrations increase when the SO₂ concentration increases.

We added the sentence below at page 12 line 13 to clarify.

Given the consistency between sulfates and organosulfates concentrations as shown in Figure S6 and S7, it is most likely that as the amount of sulfate increase with increasing initial SO₂ concentration, the concentration of organosulfates will also increase.

4. The authors concluded that the competitive reactions of OH with SO₂ and cyclohexene is the main reason for the change of SOA yield under 40 ppb of SO₂ initial concentration. However, this statement is not supported by any kinetic data or estimates based on published rate constants and measured reactant/product concentrations. Same for the reactions for the reaction of O₃+cyclohexene, what are the lifetimes of cyclohexene against OH versus O₃ oxidation? The authors should at least provide some estimates before reaching the conclusion of competitive reactions, as well as completely ignoring the potential contribution of SOA generated from O₃+cyclohexene in this study.

Although about 40% of total cyclohexene was estimated to react with O₃, the proportion of SOA formed from this reaction was not clear. From Sarrafzadeh's study,

the SOA yield was considered as a function of OH concentration (Sarrafzadeh et al., 2016). Compared to Sarrafzadeh's study, a lower O₃ concentration was used in our study, and hence the contribution of O₃ reaction to SOA might be less important. In addition, from available rate constants, the lifetime of cyclohexene based on the reaction with O₃ (5.5 h) is more than twice higher than the lifetime based on the reaction with OH (2.5 h). This further indicates that the reaction with O₃ might be less important.

For clarification, we added the following at page 9 line 8.

The rate constants of O₃ + cyclohexene and OH + cyclohexene reactions were determined to be 7.44×10^{-17} and 6.09×10^{-11} cm³ molecule⁻¹ s⁻¹, respectively, corresponding to 5.5 h and 2.5 h lifetimes for cyclohexene (Treacy et al., 1997; Rogers, 1989). Hence, it is likely that the cyclohexene reaction with O₃ would be less important than the reaction with OH in our study.

The text below was added at page 9 line 23.

The rate constant for the OH + SO₂ reaction was estimated to be 2.01×10^{-12} cm³ molecule⁻¹ s⁻¹, corresponding to a SO₂ lifetime of 69 h (Atkinson et al., 1997). This reaction is much slower than the cyclohexene + OH reaction, suggesting that OH + SO₂ reaction has very little impact on the OH concentration in the system. In our experiment, the decrease in the SOA yield with SO₂ addition might then not be attributed to its reaction with OH. It is also possible that the SO₂ addition could change the chemistry of the photooxidation process and suppress the oxygenation of products (Friedman et al., 2016; Liu et al., 2015). Comparing the HR-MS results at different initial SO₂ concentrations, the proportion of low molecular weight components increases with increasing SO₂ concentration. Molecular weights have negative correlation with volatility, which could also make the SOA yield to decrease.

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Treacy, J., Curley, M., Wenger, J., and Sidebottom, H.: Determination of Arrhenius parameters for the reactions of ozone with cycloalkenes, *J. Chem. Soc. Faraday. T.*, 93, 2877-2881, doi: 10.1039/A701794d, 1997.

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Report #2

The authors have tried to improve the quality of the manuscript, but unfortunately some of the changes raise additional questions. Generally the quality of citations and figures have improved, but unfortunately new problems with correct citation of literature has developed.

I will try to limit my comments to the responses to my previous questions.

Page 5 line 25. The authors have now stated the detection limit:

"For the IC analysis, the limit of detection was 0.2 mg mL⁻¹. "

I assume that mg should be corrected to microgram. Is this assumption correct?

The corresponding air concentration would be:

$0.2 \mu\text{g/mL} * 7 \text{ mL extraction volume} = 1.4 \mu\text{g}$.

The sampling air volume was 0.3 m³, and the detection limit was thus $1.4 \mu\text{g}/0.3 \text{ m}^3 = 4.7 \mu\text{g}/\text{m}^3$.

So the detection limit is 4.7 microgram/m³, which is around the concentration for many measurements shown in Figure 3.

The authors must explain and discuss the uncertainty of the IC analyses. If my calculation above is correct, the IC analyses are highly uncertain.

We made a mistake here. We have checked and updated the detection limit (0.005 $\mu\text{g mL}^{-1}$).

Page 6 line 31: "This is in agreement with the finding that wood soot, a minor source of SO₂ (Reddy and Venkataraman, 2002), resulted in a measurable positive deviation to the VOCs/NO_x photooxidation reaction system without background aerosol (Jang et al., 2002)."

The authors have tried to make the sentence clearer, but not yet succeeded. First of all it is not clear how wood soot can be a source of SO₂ gas, but the whole sentence makes more sense when it is compared with the original sentences from the citation.

Jang et al., 2002 wrote: "Compared to fossil fuels, biomass combustion is a minor source of SO₂ (28). Yet wood soot also resulted in a measurable positive deviation from theoretical partitioning compared to the photooxidation reaction system of α -pinene and NO_x without background aerosol."

Here it is clear that it is not wood soot, but biomass combustion, which is a minor source of SO₂ in the environment in general.

In the present context it would be much more relevant to cite studies on new particle

formation from SO₂/H₂SO₄.

We have modified this sentence to be:

It is evident from Figure 1 that even small amounts of SO₂ affect the new particle formation substantially, as observed in previous studies (Chu et al., 2016; Liu et al., 2016).

Page 7 line 20:

Before: "Condensation onto existing aerosol particles was prior to the occurrence of new particle formation."

Changed to: "New particles were formed by vapor condensation onto existing aerosol particles."

The new sentence does not make sense. New particles are never formed by growth of existing particles! After reading the text several times I now understand that you mean that condensation was favored compared to new particle formation in the system. Please rewrite to explain your statement correctly.

We have rewritten this sentence.

As long as there was enough seed particle surface area, vapor condensation onto existing aerosol particles was favored compared to the formation of new particles, and this condensation would be the main contribution to the increase of SOA mass.

Page 9 line 20. The authors have added this sentence: "It was also found that during the Investigation of Sulfur Chemistry in the Antarctic Troposphere, the OH tends to increase when the influx of SO₂ from above decreases (Mauldin et al., 2004), which means that there is a negative correlation between OH and SO₂ in real atmosphere."

It not clear how a study in Antarctica has any relevance to the conditions of the present laboratory study on cyclohexene and SO₂ in a high NO_x environment.

We agree that this statement does not directly address the issue raised in the previous report. Hence, we deleted it and solely consider our previous first suggested response to this comment, which is below:

Moreover, in real atmospheric situations where O₃ is found in much higher proportion than OH, cyclohexene would mainly react with O₃ to produce Criegee intermediates, which are good SO₂ oxidizers, and significantly less SOA than in the chamber will be formed.

Page 12 line 6-13: My previous question regarding composition and response between samples was only partly answered, so I will try to explain it more specifically: Did the "concentration" (or signal intensity) of organosulfates increase with SO₂ concentration?

Currently the authors show a normalized mass spectrum, but I would like to know the variation in intensity.

More HR-MS analyses were performed, with 0 ppb and 236 ppb initial SO₂ concentrations (see Figure S6). As can be seen from this figure, the SOA composition changes with initial SO₂ concentrations. The NL of the HR-MS results is 2.71E5, 2.28E5 and 1.78E5 when the SO₂ concentration is 236, 94 and 0 ppb, respectively. Since a liquid chromatography for quantitative analysis was not combined with the HR-MS in our study, all the HR-MS results were obtained through a direct injection. Combined with the FTIR spectra results, we believe the amount of organosulfates in SOA was increased with initial SO₂ concentration.

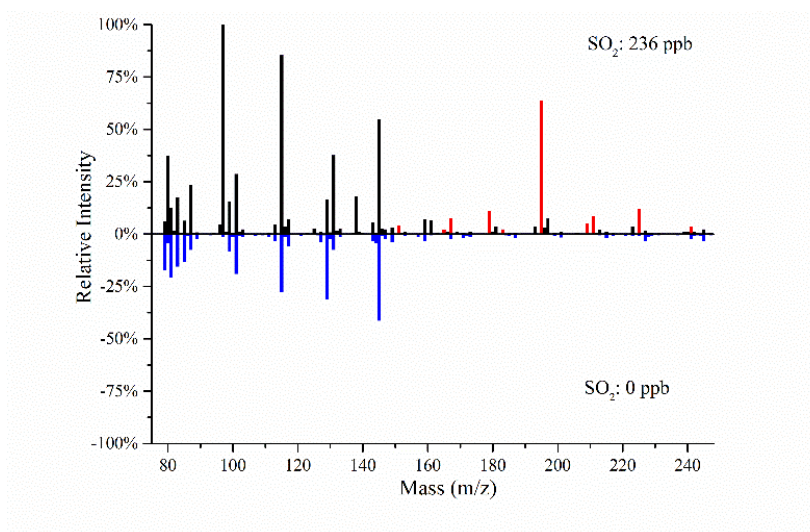


Figure S6: Comparison of SOA ESI-HR-MS spectra with different initial SO₂ concentrations.

We added the following at page 12 line 13.

Comparing HR-MS data when initial concentrations were 0 ppb and 236 ppb reveals that the bands representing organosulfates do not appear 0 ppb of SO₂. Peaks at m/z larger than 150 were undetectable at initial SO₂ concentration of 0 ppb, while products without sulfur peaked at both concentrations, with the only difference being their intensities. This implies that the process of SOA formation strongly depends on initial SO₂ concentrations.

Additional comments:

Please correct the figure caption for Figure 3 to describe the current version.

We have fixed it.

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Chu, B. W., Zhang, X., Liu, Y. C., He, H., Sun, Y., Jiang, J. K., Li, J. H., and Hao, J. M.: Synergetic formation of secondary inorganic and organic aerosol: effect of SO₂ and NH₃ on particle formation and growth, *Atmos. Chem. Phys.*, 16, 14219-14230, doi: 10.5194/acp-16-14219-2016, 2016.

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Report #3

In this study the authors report the formation of organosulfates from the photooxidation of cyclohexene. The manuscript went through the ACPD open discussion process and got mixed reviews, with a number of good suggestions from the more critical reviewers. Based on my examination of the authors' responses to the reviewers' comments, I believe the concerns of the reviewers have been mainly addressed. However, the authors should (1) better discuss the oxidation regime in their experiments and (2) fix the different errors present in the revised version.

To solve the discussion about the impact of ozone in the oxidation of cyclohexene, the authors need to run a simple box model to estimate the degradation pathways. As they are, the explanations are not convincing. Indeed, the authors claim the ozonolysis doesn't contribute to depletion of cyclohexene but the authors didn't propose any estimation of the concentration of OH radicals. It is also surprising that the addition of large quantities of SO₂ does not impact more the decay of cyclohexene (e.g. no difference between 40 and 90 ppb).

Finally that's not true that the ozonolysis of cyclohexene produces sCI (c.f. Donahue et al., 2011).

The paragraph 3.2 should be rewritten to discuss the results from a more objective scientific points.

We estimate the degradation pathways of cyclohexene according to the changes of O₃ and cyclohexene concentrations, and the rate constants of both O₃+cyclohexene and OH·+cyclohexene reactions. About 60% of the total cyclohexene was oxidized by OH, while another fraction of cyclohexene oxidation is attributed to the reaction with O₃. The addition of large quantities of SO₂ does not lead to a remarkable cyclohexene decrease. This might be due to the fact that the SO₂+OH· reaction quickly reaches saturation, so that additional SO₂ does not lead an obvious increase in the consumption of OH. As shown in Figure 1, the particle number concentrations were practically maintained steady when the SO₂ concentrations were varied systematically between 30 and 105 ppb. Nucleation was strongly dependent on the abundance of H₂SO₄, which was based on the oxidized SO₂ + OH radicals (Sihto et al., 2006; Xiao et al., 2015). This feature may indicate that no more sulfates were formed when SO₂ concentration was in large excess (>30 ppb) and the OH radicals being insufficient. Higher initial SO₂ concentrations may neither decrease the OH concentration substantially, nor induce important cyclohexene decay.

Concerning the sCI formation from cyclohexene-ozone system, Donahue et al.

pointed out that “although we see no evidence of sCI formation from cyclohexene ozonolysis, secondary ozonides have been reported from this system, suggesting sCI formation” (Drozd and Donahue, 2011). The quantitative measure of the sCI yield from cyclohexene ozonolysis was found to be very low, about 3% (Hatakeyama et al., 1984). This is because the Criegee intermediates from endocyclic alkenes are formed with more energy release than those from exocyclic alkenes and hence, are less likely to be stabilized (Chuong et al., 2004). We have revised the discussion about sCI.

The revised part in paragraph 3.2 is as follows.

In the presence of SO₂, enhanced SOA formation was attributed to acid-catalyzed heterogeneous reactions (Jang et al., 2002; Xu et al., 2014). When studying the effect on acidic seed on the growth of isoprene- and α -pinene-based SOA, it was shown that FTIR peaks at 1180, 1050 cm⁻¹ are indicators of acid-catalyzed heterogeneous reactions since these peaks could not otherwise be observed in non-acidic conditions (Czoschke et al., 2003 and references therein). In our experiment, two similar bands located at 1195 and 1040 cm⁻¹, were observed (see Figure S4), which supports the presence of acid-catalyzed reactions in our experiment. These observed peaks are prominent in IR spectra from SOA formation in an acidic particle environment (Jang et al., 2002; Czoschke et al., 2003). In our study, the intensities of these peaks were very weak when initial SO₂ concentrations were lower than 44 ppb, indicating that acid-catalyzed reactions are not facilitated at these conditions.

However, there were some undiscovered processes that could inhibit the formation of SOA in the cyclohexene/NO_x/SO₂ system. The competitive reaction between SO₂ and cyclohexene might be one of the reasons for the decrease in the SOA yield. For example, SO₂ could be oxidized by OH to form H₂SO₄ (Somnitz, 2004). Due to the presence of O₃ in our system, the formation of Criegee intermediates and their reactions with SO₂ could be expected (Criegee, 1975). The rate constants of O₃ + cyclohexene and OH + cyclohexene reactions were determined to be 7.44×10⁻¹⁷ and 6.09×10⁻¹¹ cm³ molecule⁻¹ s⁻¹, respectively, corresponding to 5.5 h and 2.5 h lifetimes for (Treacy et al., 1997; Rogers, 1989). Hence, it is likely that the cyclohexene reaction with O₃ would be less important than the reaction with OH in our study. However, the importance of SO₂+sCI reactions could be limited due to the kinetics and low sCI yield (Stewart et al., 2013; Keywood et al., 2004; Hatakeyama et al., 1984). As mentioned above, the photooxidation in this study was at high-NO_x conditions and the OH was the main limiting factor for SOA formation because of its relatively low concentration. The change of cyclohexene concentration with time at different initial SO₂ concentrations is shown in Figure S5, wherefrom it can be seen that the reacted cyclohexene concentration at 0 ppb initial SO₂ concentration was slightly higher than that at 90 ppb. The consuming rate of cyclohexene was higher without SO₂ in the chamber, which means that if there was a competition reaction, its effect was very limited. At lower OH concentration condition caused by the reaction between SO₂ and OH, the formation of SOA was inhibited.

The rate constant for the OH + SO₂ reaction was estimated to be 2.01×10⁻¹² cm³ molecule⁻¹ s⁻¹, corresponding to a SO₂ lifetime of 69 h (Atkinson et al., 1997). This

reaction is much slower than the cyclohexene + OH reaction, suggesting that OH + SO₂ reaction has very little impact on the OH concentration in the system. In our experiment, the decrease in the SOA yield with SO₂ addition might then not be attributed to its reaction with OH. It is also possible that the SO₂ addition could change the chemistry of the photooxidation process and suppress the oxygenation of products (Friedman et al., 2016; Liu et al., 2015). Comparing the HR-MS results at different initial SO₂ concentrations, the proportion of low molecular weight components increases with increasing SO₂ concentration. Molecular weights have negative correlation with volatility, which could also make the SOA yield to decrease. Moreover, in real atmospheric situations where O₃ is found in much higher proportion than OH, cyclohexene would mainly react with O₃ to produce Criegee intermediates, which are good SO₂ oxidizers, and significantly less SOA than in the chamber will be formed. Accordingly, SOA yield showed descending trend with the increase of SO₂ concentration when the initial SO₂ concentration was lower than 40 ppb.

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Photooxidation of cyclohexene in the presence of SO₂: SOA yield and chemical composition

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Abstract. Secondary organic aerosol (SOA) formation from cyclohexene/NO_x system with various SO₂ concentrations under UV light was investigated to study the effects of cyclic alkenes on the atmospheric environment in polluted urban areas. A clear decrease at first and then increase of the SOA yield was found with increasing SO₂ concentrations. The lowest SOA yield was obtained when initial SO₂ concentration was in the range of 30-40 ppb, while higher SOA yield compared to that without SO₂ could not be obtained until the initial SO₂ concentration was higher than 85 ppb. The decreasing SOA yield might be due to the fact that the promoting effect of acid-catalyzed reactions on SOA formation was less important than the inhibiting effect of decreasing OH concentration at low initial SO₂ concentrations, caused by the competition reactions of OH with SO₂ and cyclohexene. The competitive reaction was an important factor for SOA yield and it should not be neglected in photooxidation reactions. The composition of organic compounds in SOA was measured using several complementary techniques including Fourier transform infrared (FTIR) spectrometer, ion chromatograph (IC) and electrospray ionization high-resolution quadrupole mass spectrometer (ESI-HR-MS). We present the first evidence that organosulfates were produced from the photooxidation of cyclohexene in the presence of SO₂.

1 Introduction

Alkenes are widely emitted from biogenic and anthropogenic sources (Kesselmeier et al., 2002; Chin and Batterman, 2012), and their gas-phase oxidation reactions with OH, NO₃, or O₃ are among the most important processes in the atmosphere (Atkinson, 1997; Stewart et al., 2013; Paulson et al., 1999). The reactions of ozone with alkenes are an important source of free radicals in the lower atmosphere, which influence the oxidative capacity of the atmosphere (Paulson and Orlando, 1996). Some of the products have sufficiently low vapor pressures to condense with other gaseous species, and contribute to the secondary organic aerosol (SOA) mass (Sarwar and Corsi, 2007; Sakamoto et al., 2013; Nah et al., 2016; Kroll and Seinfeld, 2008; Hallquist et al., 2009). SOA formation from VOCs oxidation has been receiving significant attention since recent years due to its large implication in the formation of atmospheric fine particulate matter (Jimenez et al., 2009). SOA has significant

impacts on human health (Pope III and Dockery, 2006), air quality (Kanakidou et al., 2005; Jaoui et al., 2012; McFiggans et al., 2006), and global climate change (Hansen and Sato, 2001; Adams et al., 2001; Pokhrel et al., 2016).

Although cyclic alkenes widely exist in the atmosphere, their gas-phase oxidation has received less attention than that of linear or branched alkenes (Sipilä et al., 2013). Cyclohexene is an important industrial chemical (Sun et al., 2013), and is also widespread in urban areas (Grosjean et al., 1978). Cyclohexene has been extensively studied as a monoterpene surrogate for inferring oxidation mechanisms and aerosol formation characteristics, because it has the basic structural unit in abundant biogenic monoterpenes and sesquiterpenes (Carlsson et al., 2012; Keywood et al., 2004b). The rate constants for gas-phase reactions of cyclohexene with OH, O₃ and NO₃ were measured at room temperature to be $(6.4 \pm 0.1) \times 10^{-11}$, $(8.1 \pm 1.8) \times 10^{-17}$ and $(5.4 \pm 0.2) \times 10^{-13}$ cm³ molecule⁻¹ s⁻¹, respectively (Stewart et al., 2013; Aschmann et al., 2012), and a correlation between the logarithm of the rate constants and the molecular orbital energies for simple cyclic alkenes was observed. The effect of pressure and that of the presence of SO₂ on the formation of stable gas-phase products and SOA from ozonolysis of cyclohexene were investigated (Carlsson et al., 2012). It was found that the collisional stabilization of initial clusters was an important aspect for SOA formation processes involving sulfuric acid (H₂SO₄) and organic compounds. The effect of the structure of the hydrocarbon parent molecule on SOA formation was investigated for a series of cyclic alkenes and related compounds (Keywood et al., 2004b), and the SOA yield was found to be a function of the number of carbons present in the cyclic alkenes ring. The relative SOA yields from ozonolysis of cyclic alkenes can be quantitatively predicted from properties of the parent hydrocarbons, like the presence of a methyl group and an exocyclic double bond.

SO₂, one of the most important inorganic pollutants in urban area, plays an important role on SOA formation (Wang et al., 2005; Lonsdale et al., 2012; Liu et al., 2016). Seasonal variations of SO₂ concentrations were found to be consistent with seasonal variations of PM_{2.5} concentration (Cheng et al., 2015). Smog chamber simulations have indicated that SO₂ could enhance the formation of SOA from VOCs oxidation under acidic conditions by increasing aerosol acidity and ammonium sulfate aerosol formation (Edney et al., 2005; Liu et al., 2016; Attwood et al., 2014). Anthropogenic SO₂ emissions can impact new particle formation, and SOA composition (Lonsdale et al., 2012).

Despite the existence of organosulfates in ambient aerosols was first observed in 2005 (Romero and Oehme, 2005), proper identification of these aerosols was made two years later. In a series of chamber experiments studies, it was shown that organosulfates present in ambient aerosols collected from various locations mostly originate from acid-catalyzed reactions of SOA formed from photooxidation of α -pinene and isoprene (Surratt et al., 2007). Recently, different kinds of organosulfates have been observed in SOA around the world, and organosulfates have been identified as a group of compounds that have an important contribution to the total amount of SOA in the atmosphere (Surratt et al., 2008; Froyd et al., 2010; Kristensen and Glasius, 2011; Tolocka and Turpin, 2012; Wang et al., 2015). Laboratory chamber studies showed that OH/NO_x/O₃-initiated reactions of BVOCs, such as isoprene, α -pinene, β -pinene, and limonene with sulfates or sulfuric acid are the main processes for organosulfates formation (Surratt et al., 2007; Surratt et al., 2008; Hatch et al., 2011). Despite the well-recognized presence of organosulfates in SOA, their formation and transformation processes can be complex and varied, depending on the nature of the original organic compound involved. Extensive studies on their formation have been performed and several

mechanisms based on a variety of reactions have been proposed. Using nuclear resonance techniques, isoprene-derived epoxides formed during isoprene photooxidation reactions were found to be important intermediates for organonitrates and organosulfates formation via potential SOA reactions (Darer et al., 2011; Hu et al., 2011). The authors further found that organonitrates could easily be transformed to organosulfates during hydrolysis in the presence of sulfate. Some studies also showed that 2-methyl-3-buten-2-ol (MBO), due to its larger emissions than isoprene in some regions (Baker et al., 1999), is an important precursor for organosulfates and SOA in the atmosphere, through its reactions with OH under NO and aerosol acidity conditions, and from acid-catalyzed reactive uptake of MBO-based epoxides formed during MBO photooxidation (Mael et al., 2015; Zhang et al., 2012; Zhang et al., 2014). Organosulfates formation was also found from oxidation of hydroxyhydroperoxides (Riva et al., 2016) and from heterogeneous reactions of SO₂ with selected long-chain alkenes and unsaturated fatty acids (Passananti et al., 2016).

Reactions with sulfates or H₂SO₄ were the main formation processes of organosulfates. Qualitative analyses of the organosulfates in SOA have been gaining more attention and development since recent years (Lin et al., 2012; Shalamzari et al., 2013; Staudt et al., 2014). Riva et al. investigated the formation of organosulfates from photooxidation of polycyclic aromatic hydrocarbons and found that, in the presence of sulfate aerosol, this photooxidation was a hitherto unrecognized source of anthropogenic secondary organosulfur compounds (Riva et al., 2015). A more complete structural characterization of polar organosulfates that originate from isoprene SOA was performed (Shalamzari et al., 2013), and an organosulfate related to methyl vinyl ketone and minor polar organosulfates related to croton aldehyde were identified. However, there was no report about the yield and chemical composition of SOA obtained from photooxidation of cyclohexene in the presence of SO₂.

In the present work, we investigated the SOA yields and chemical composition during cyclohexene photooxidation under different SO₂ concentrations conditions. A better understanding of the magnitude and chemical composition of SOA from different SO₂ concentrations will contribute to more accurate SOA prediction from anthropogenic sources and give valuable information related to air pollution in urban environments.

2 Methods

2.1 Chamber description

The experiments were performed in a 400 L Teflon FEP film chamber (wall thickness 125 μm) at the Institute of Atmospheric Physics, Chinese Academy of Sciences, Beijing. The details of the chamber, including the experimental setup and analysis techniques have been described elsewhere (Du et al., 2007; Jia and Xu, 2014), and only a brief description is presented here. The reactor was surrounded by 12 black light lamps (GE F40BLB) with emission band centered at 365 nm, which were used to simulate the spectrum of the UV band in solar irradiation. Stainless steel was covered on the chamber interior walls to maximize and homogenize the interior light intensity. The effective light intensity near the ultraviolet region plays a decisive role in the formation of photochemical smog (Presto et al., 2005a). The effective light intensity of the

chamber was represented by the photolysis rate constant of NO_2 . In our study, the average effective light intensity was determined to be 0.177 min^{-1} . Both inlet and outlet of the chamber were made of Teflon material. Atmospheric pressure was maintained in the chamber at all times. All the experiments were performed at room temperature ($307 \pm 2 \text{ K}$) under dry conditions ($\text{RH} < 10\%$). The wall loss is the decrease of the concentration of reactive gas phase species caused by adsorption on the inner wall of the reactor. Residual reactant and product on the inner wall can also react with the gas phase species, which is another important reason for wall loss. The wall loss can directly affect the quantitative evaluation of the photooxidation rate and SOA yield. A correct estimation of the wall loss is therefore necessary in a reliable analysis of the experimental results of SOA yield. In the present study, the wall loss of cyclohexene in the chamber could be neglected since no decrease in its concentration was observed. The wall loss of O_3 , NO_x , and SO_2 were first order because $\ln([X]_0/[X]_t)$ had a good correlation with time ($R^2=0.994, 0.944, 0.999$ for O_3 , NO_x , and SO_2 , respectively). The measured wall loss rate constants for O_3 , NO_x and SO_2 were 5.05×10^{-6} , 7.04×10^{-6} and $6.39 \times 10^{-6} \text{ s}^{-1}$, respectively. The average value of the wall loss rate constant of particles was $4.7 \times 10^{-5} \text{ s}^{-1}$, and the measured particle concentrations in this study were corrected using the same method as Pathak et al. (Pathak et al., 2007). Typical profiles of the gas and particle phases are given in Figure S1 of the Supplementary material.

Prior to each experiment, the chamber was cleaned by purging with purified dry air for at least 8 h until residual hydrocarbons, O_3 , NO_x or particles could not be detected in the reactor. Known amounts of cyclohexene was injected into a 0.635 cm diameter Teflon FEP tube and dispensed into the chamber by purified dry air. Typical initial cyclohexene concentrations were 500 ppb. NO_x was injected by a gas-tight syringe to make the mixing ratio of NO_x in the reactor around 95 ppb during all the experiments. The mixed concentration ratios of cyclohexene/ NO_x were in the range 4.4-6.9. SOA formation experiments were carried out under UV irradiation in the presence of NO_x to produce O_3 and OH radicals for cyclohexene oxidation. Although initial VOCs, NO_x and average OH concentrations were different from typical urban conditions, efforts were made to maintain the initial concentrations of the reactants as similar as possible to make sure the effect of SO_2 was the main reason for the changes in the SOA yield. More details on the experimental conditions are shown in Table 1.

2.2 Gas and particle measurements

Ozone concentration in the reactor was measured using ozone analyzer (Model 49C, Thermo Electron Corporation, USA). A $\text{NO-NO}_2\text{-NO}_x$ analyzer (Model 42C, Thermo Electron Corporation, USA) was used to monitor the NO_x concentration. Measurement of SO_2 concentration was made using a SO_2 analyzer (Model 43i-TLE, Thermo Electron Corporation, USA). The uncertainty of the O_3 , NO_x and SO_2 measurement was less than $\pm 1\%$. The detection limits of the different monitors were 0.40 ppb, 0.50 ppb, and 0.05 ppb for NO_x , O_3 , and SO_2 , respectively.

Two Tenax absorption tubes (150 mm length \times 6 mm O.D., 0.2 g sorbent) were used to collect the sample before the UV lights were turned on and at the end of each experiment, respectively. The volume of the sample was 60 mL and the sampling time was 3 min. The concentrations of cyclohexene were analyzed by thermal desorption-gas chromatography-

mass spectrometry (TD-GC-MS). A thermal desorption unit (Master TD, Dani, Italy) was combined with a 6890A gas chromatograph (6890A, Agilent Tech., USA) interfaced to a 5975C mass selective detector (5975C, Agilent Tech., USA). The GC was equipped with a HP-5MS capillary column (30 m × 0.25 mm, 0.25 μm film thickness). The TD temperature was 280 °C, and the sampling time was 3 min. The GC-MS temperature program was as follows: the initial temperature of 40 °C was held for 4 min, and then raised to 300 °C at a rate of 20 °C min⁻¹. The inlet temperature was set at 250 °C and the transfer line at 200 °C. The ionization method in MS was electron impact ionization, and helium was used as the carrier gas at a constant flow (1.2 mL min⁻¹). Because a very diverse range of compounds might be present in the samples, the SCAN mode (36-500 amu) was used in the MS detector. This mode is known to be a classical and typical detection method for GC-MS analysis. The results were analyzed with MSD Productivity ChemStation.

Particle number concentrations and size distributions were measured with a scanning mobility particle sizer (SMPS), which consists of a differential mobility analyzer (DMA model 3081, TSI Inc., USA) and a condensation particle counter (CPC model 3776, TSI Inc., USA). A sheath flow/aerosol flow relationship of 3.0/0.3 L min⁻¹ was used for the measurements. The particle size was measured in the range of 14 to 710 nm, and each scan was 180 s. An aerosol density of 1.2 g cm⁻³ was assumed to convert the particle volume concentration into the mass concentration (Zhang et al., 2015). Size distribution data were recorded and analyzed using the TSI AIM software v9.0.

2.3 Products composition analysis

The chemical composition of SOA was important for analyzing the degree of cyclohexene oxidation, and it was used to evaluate the transformation from gas phase to particle phase. Particle phase chemical composition was studied by means of Fourier transform infrared (FTIR) spectrometer (Nicolet iS10, Thermo Fisher, USA). The aerosols were sampled through a Dekati low pressure impactor (DLPI, DeKati Ltd, Finland). The impactor was connected to a pump working at a flow rate of 10 L min⁻¹ while sampling a total volume of 300 L of gas for each experimental run. Aerosols, from 108 to 650 nm diameter, were collected on an ungreased zinc selenide (ZnSe) disk (25 mm in diameter) for FTIR measurements.

The characteristic bands of inorganic and organic sulfates overlap in the IR spectrum. In order to distinguish between the inorganic and organic sulfates, ion chromatograph (IC, Dionex ICS-900, Thermo Fisher, USA) was used to analyze the inorganic sulfate anion (SO₄²⁻) in SOA. For the IC analysis, the limit of detection was 0.005 mg mL⁻¹. SOA collected on ZnSe disks was firstly dissolved in high purity water (7 mL) and then measured by IC for SO₄²⁻ concentrations. The anions were analyzed with a Dionex IonPac AS14A analytical column and an anion self-regenerating suppressor Dionex ASRS was used as eluent. The flow rate was 1.0 mL min⁻¹ with a mixture of 8.0 mmol L⁻¹ Na₂CO₃ and 1.0 mmol L⁻¹ NaHCO₃ for anions analyses. The suppressing current was 50 mA.

Chemical characterization of aerosols from photooxidation of cyclohexene was performed using an electrospray ionization high-resolution quadrupole mass spectrometer (ESI-HR-MS, Thermo Fisher, USA) operated in negative (-) ion mode, which was calibrated using the manufacturer's calibration standards mixture allowing for mass accuracies <5 ppm in external calibration mode. The ionization voltage was 4.2 kV and the capillary temperature was set at 300 °C. N₂ was used as both

the sheath gas (70 U) and auxiliary gas (30 U). SOA was collected on the aluminium foil using the same method as FTIR analysis and then extracted with 1 mL of acetonitrile. The aluminium foil was used due to its ease to handle and its non-reactivity with the sample. A total volume of 300 L was sampled at a flow rate of 10 L min⁻¹. A volume of 5 μL of the extraction and a direct injection were used for the measurement. Xcalibur 2.2 software (Thermo Fisher, USA) was used for the calculation of chemical formula from the accurate measurement of m/z values.

2.4 Chemicals

The chemicals used and their stated purities were as follows: cyclohexene (99%) was obtained from Aldrich and used without further purification. A zero air generator (Model 111, Thermo Scientific, USA) was used to generate clean air. The zero air has no detectable non-methane hydrocarbons (NMHC < 1 ppb), NO_x (< 1 ppb), low O₃ concentration (< 3 ppb), and low particle numbers (< 5 cm⁻³), and relative humidity (RH) below 10%. Ozone was produced from O₂ via electrical discharge using a dynamic gas calibrator (Model 146i, Thermo Scientific, USA). NO₂ (510 ppm), NO (50 ppm) and SO₂ (25 ppm) with ultra-pure N₂ (99.999%) as background gas was purchased from Beijing Huangyuan Gas Co., Ltd., China.

3 Results and discussion

3.1 Effect of SO₂ on SOA number concentrations

The particle number concentrations at the maximum SOA yield for the cyclohexene/NO_x/SO₂ system with different initial SO₂ concentrations are shown in Figure 1. The cooperation of the maximum number concentration and the particle number concentrations at the maximum SOA yield is shown in Figure S2. After the black light lamps were turned on, the SOA number concentrations increased rapidly to reach the maximum within 0.5 h in each experiment. Subsequently, the particle number concentrations gradually decreased accompanied by the growth of particle size by coagulation. The SOA mass concentration kept increasing until its maximum was reached (after ~2 h). Both types of particle number concentrations had similar trends against initial SO₂ concentrations. In general, maximum particle number concentrations were three times higher than the particle number concentrations at the maximum SOA yield. In the remainder of this paper, in order to better elaborate the effect of SO₂ on the formation of particles, the particle number concentration refers to the particle number concentrations at the maximum SOA yield.

The particle number concentration increased with initial SO₂ concentration, and this increase could be divided into two stages: increasing stage and stable stage. In the increasing stage, with the initial SO₂ concentration increasing from 0 ppb to 30 ppb, the particle number concentration grew significantly under low initial SO₂ concentration (<5 ppb), then the growth rate reduced gradually. In the stable stage, when the SO₂ concentrations were varied systematically between 30 and 105 ppb, particle number concentrations were practically maintained steady, and there was no further obvious growth as shown in Figure 1. For experiments with high initial SO₂ concentrations, the particle number concentrations were 10 times higher than those without SO₂, indicating enhanced new particle formation (NPF) when adding SO₂. **It is evident from Figure 1 that even**

small amounts of SO₂ affect the new particle formation substantially, as observed in previous studies (Chu et al., 2016; Liu et al., 2016).

Nucleation is a fundamental step in the atmospheric new particle formation. Nucleation of particles in the atmosphere has been observed to be strongly dependent on the abundance of H₂SO₄ (Sihto et al., 2006; Xiao et al., 2015). Normally, SO₂ was deemed to be oxidized by OH radicals to form H₂SO₄ through homogeneous reactions in gas phase (Calvert et al., 1978), or by H₂O₂ and O₃ through in-cloud processes in aqueous phase (Lelieveld and Heintzenberg, 1992). The aqueous phase formation of H₂SO₄ is negligible in this study (RH<10%). As the precursor of H₂SO₄, SO₂ at high concentrations would lead to more H₂SO₄ formation, and thereby increase the nucleation rates and total particle number concentrations (Sipilä et al., 2010). Because of the similar initial conditions for each experiment except SO₂, the amount of OH radicals produced was assumed to be almost equal. In the presence of high concentrations of SO₂, new particle formation was not enhanced. This feature may indicate that no more sulfates were formed when SO₂ was in large excess (>30 ppb) and the OH radicals were insufficient. The quantity of OH radicals is the main restraint on H₂SO₄ formation at high initial SO₂ concentrations, which could not lead to more H₂SO₄ formation in the present study. Therefore, the particle number concentration was maintained steady, and it was independent of the SO₂ concentrations in the second stage.

Besides, the mean diameter of particles increased with photooxidation reaction time, which suggests that not many new particles were generated after a burst increase at the initial stage of SOA formation. Once new particles are formed, there is a competition between growth of existing particles by uptake of the precursors and formation of new particles. Our result agrees with previous studies that there was no obvious increase in aerosol number concentration when additional VOCs were injected, but a significant increase in SOA mass concentration (Presto et al., 2005a). **As long as there was enough seed particle surface area, vapor condensation onto existing aerosol particles was favored compared to the formation of new particles, and this condensation would be the main contribution to the increase of SOA mass.**

3.2 Effect of SO₂ on SOA yields

SOA yield (Y) is defined as $Y = \Delta M_0 / \Delta HC$, where ΔM_0 is the produced organic aerosol mass concentration ($\mu\text{g m}^{-3}$), and ΔHC is the mass concentration of reacted cyclohexene ($\mu\text{g m}^{-3}$). The SOA yields of cyclohexene at different SO₂ concentration as determined by SMPS are shown in Figure 2. The numerical values of the aerosol mass concentration and SOA yields at different conditions are shown in Table 1.

The SOA yields in the absence of SO₂ were in the range of 2.7-3.4%, which were an order of magnitude lower than those reported in previous studies (Warren et al., 2009; Keywood et al., 2004b; Kalberer et al., 2000). There are three possible explanations to this phenomenon. (1) SOA formation is closely related to the oxidation capacity in the photooxidation experiments and, therefore, is affected by the ratio of $[\text{VOC}]_0 / [\text{NOx}]_0$ (Pandis et al., 1991). Experiments performed with different SO₂ concentrations indicate that the SOA formation is partly controlled by the ability of the system to oxidize cyclohexene and contribute to the particle mass. As indicated in Figure S3 of the Supplementary material, even at 0 ppb of SO₂, the mass concentration of SOA quickly reaches its maximum. Experiments with higher NO_x levels have been proved to

get considerably lower SOA yield than those with lower NO_x levels at the same VOCs concentration (Song et al., 2005). The reactions of organo-peroxy radicals (RO₂) with NO and NO₂ instead of peroxy radicals (RO₂ or HO₂) under high NO_x conditions resulted in the formation of volatile organic products and a decreased SOA yield (Lane et al., 2008). It was reported that SOA yield was constant for $[\text{VOC}]_0/[\text{NO}_x]_0 > 15$, but decreased considerably (by a factor of more than 4) as $[\text{VOC}]_0/[\text{NO}_x]_0$ decreased (Presto et al., 2005b). In our study, the $[\text{VOC}]_0/[\text{NO}_x]_0$ ratio was maintained at about 4.4 to 6.9. Recently, the NO_x dependence of SOA formation from photooxidation of β -pinene was comprehensively investigated (Sarrafzadeh et al., 2016), and it was shown that the NO_x-induced OH concentration was the major factor influencing the SOA yield. The impacts of NO_x on SOA formation were only moderate if the impact of NO_x on OH concentration was eliminated. OH concentration in our study was relatively insufficient, which was the main limiting factor for SOA formation.

(2) UV light is another factor influencing the SOA yield. SOA yields between dark and UV-illuminated conditions were reported to be different (Presto et al., 2005a). Exposure to UV light could reduce SOA yield by 20-40%, while more volatile products were formed (Griffin et al., 1999). (3) The temperature may have a pronounced influence on SOA yield (Qi et al., 2010; Emanuelsson et al., 2013). At low temperatures, semi-volatile organic compounds would favor the condensation of gas phase species and a higher SOA yield could be expected. Raising the chamber temperature by 10 K should cause a decrease of 10% in aerosol yield (Pathak et al., 2007). SOA yields reported in the present study were obtained at a higher temperature, (307±2 K) than 298 K used in most previous studies. On the basis of the discussion above, the SOA yield from cyclohexene in our study was lower than observed in the previous studies.

SOA yields for the cyclohexene/NO_x/SO₂ system were measured for initial SO₂ mixing ratios of 0-105 ppb. Due to the error associated with the SO₂ concentrations measurement, with stronger impact on low values than on higher values, several experiments were performed at SO₂ concentrations below 40 ppb. The experimental results showed a clear decrease at first and then an increase in the SOA yield with increasing SO₂ concentrations (Figure 2). When SO₂ concentrations increased from 0 to 40.8 ppb, there was a remarkable decrease in SOA yield, dropping by about half with the increase of SO₂ concentration. For SO₂ concentrations higher than 40.8 ppb, SOA yield increased with increasing SO₂ concentration. The highest SOA yield was found to be 3.5%, and was at 104.7 ppb SO₂ concentration. The lowest SOA yield of cyclohexene photooxidation was obtained at an initial SO₂ concentration of 40 ppb. Although the SOA yield increased gradually with the initial SO₂ concentration at concentrations higher than 40 ppb, a higher SOA yield than that in the absence of SO₂ could not be obtained when the initial SO₂ concentration was lower than 85 ppb.

Both NO and NO₂ were used as NO_x for the repeated experiments in our study. Although the photooxidation reaction could not happen in the case of NO until it was oxidized to NO₂, which means that both NO- and NO₂-initiated photooxidation reactions were actually triggered by NO₂, the chemistry of SOA formation from both processes is similar. Despite the time of occurrence of the maximum SOA concentration for the experiment with NO₂ was half an hour earlier than that for the experiment with NO, the results of SOA yield were similar.

In the presence of SO₂, enhanced SOA formation was attributed to acid-catalyzed heterogeneous reactions (Jang et al., 2002; Xu et al., 2014). When studying the effect on acidic seed on the growth of isoprene- and α -pinene-based SOA, it was shown

that FTIR peaks at 1180, 1050 cm^{-1} are indicators of acid-catalyzed heterogeneous reactions since these peaks could not otherwise be observed in non-acidic conditions (Czoschke et al., 2003 and references therein). In our experiment, two similar bands located at 1195 and 1040 cm^{-1} , were observed (see Figure S4), which supports the presence of acid-catalyzed reactions in our experiment. These observed peaks are prominent in IR spectra from SOA formation in an acidic particle environment (Jang et al., 2002; Czoschke et al., 2003). In our study, the intensities of these peaks were very weak when initial SO_2 concentrations were lower than 44 ppb, indicating that acid-catalyzed reactions are not facilitated at these conditions.

However, there were some undiscovered processes that could inhibit the formation of SOA in the cyclohexene/ NO_x / SO_2 system. The competitive reaction between SO_2 and cyclohexene might be one of the reasons for the decrease in the SOA yield. For example, SO_2 could be oxidized by OH to form H_2SO_4 (Somnitz, 2004). Due to the presence of O_3 in our system, the formation of Criegee intermediates and their reactions with SO_2 could be expected (Criegee, 1975). The rate constants of $\text{O}_3 + \text{cyclohexene}$ and $\text{OH} + \text{cyclohexene}$ reactions were determined to be 7.44×10^{-17} and 6.09×10^{-11} $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$, respectively, corresponding to 5.5 h and 2.5 h lifetimes for cyclohexene (Treacy et al., 1997; Rogers, 1989). Hence, it is likely that the cyclohexene reaction with O_3 would be less important than the reaction with OH in our study. However, the importance of $\text{SO}_2 + \text{sCI}$ reactions could be limited due to the kinetics and low sCI yield (Stewart et al., 2013; Keywood et al., 2004a; Hatakeyama et al., 1984). As mentioned above, the photooxidation in this study was at high- NO_x conditions and the OH was the main limiting factor for SOA formation because of its relatively low concentration. The change of cyclohexene concentration with time at different initial SO_2 concentrations is shown in Figure S5, wherefrom it can be seen that the reacted cyclohexene concentration at 0 ppb initial SO_2 concentration was slightly higher than that at 90 ppb. The consuming rate of cyclohexene was higher without SO_2 in the chamber, which means that if there was a competition reaction, its effect was very limited. At lower OH concentration condition caused by the reaction between SO_2 and OH, the formation of SOA was inhibited.

The rate constant for the $\text{OH} + \text{SO}_2$ reaction was estimated to be 2.01×10^{-12} $\text{cm}^3 \text{ molecule}^{-1} \text{ s}^{-1}$, corresponding to a SO_2 lifetime of 69 h (Atkinson et al., 1997). This reaction is much slower than the cyclohexene + OH reaction, suggesting that $\text{OH} + \text{SO}_2$ reaction has very little impact on the OH concentration in the system. In our experiment, the decrease in the SOA yield with SO_2 addition might then not be attributed to its reaction with OH. It is also possible that the SO_2 addition could change the chemistry of the photooxidation process and suppress the oxygenation of products (Friedman et al., 2016; Liu et al., 2015). Comparing the HR-MS results at different initial SO_2 concentrations, the proportion of low molecular weight components increases with increasing SO_2 concentration. Molecular weights have negative correlation with volatility, which could also make the SOA yield to decrease. Moreover, in real atmospheric situations where O_3 is found in much higher proportion than OH, cyclohexene would mainly react with O_3 to produce Criegee intermediates, which are good SO_2 oxidizers, and significantly less SOA than in the chamber will be formed. Accordingly, SOA yield showed descending trend with the increase of SO_2 concentration when the initial SO_2 concentration was lower than 40 ppb.

When the initial SO₂ concentration was greater than 40 ppb, the acid-catalyzed heterogeneous formation of SOA became more significant (Figure 2). The same SOA yield was obtained in the absence of SO₂ and at 85 ppb initial SO₂ concentration. The competitive reaction plays an important role on SOA formation, and it should be taken into account in SOA simulation models or air quality models for more accurate prediction. The acid-catalyzed reactions gradually increased with initial SO₂ concentration for SOA yield increase. The formation of low volatile organics (e.g. organosulfates) by photooxidation in the presence of SO₂ might be another reason for the increase of the SOA yield.

3.3 Organosulfates formation

When SO₂ was added into the chamber, the acidic aerosol particles were formed by photooxidation of SO₂ in a reaction initiated by OH. The amount of SO₄²⁻ in particle phase and the consumption of SO₂ (Δ SO₂) with different initial SO₂ concentrations are shown in Figure 3. The changes with initial SO₂ concentrations were not uniform between the SO₄²⁻ concentration and Δ SO₂, which indicates that besides SO₄²⁻, other products were formed from the reaction of SO₂. Typical IR spectra of aerosols from the cyclohexene/NO_x/SO₂ system under different SO₂ concentrations are presented in Figure 4. Based on the peak positions in the IR spectra, the functional groups represented by each peak are summarized as follows: the broadband at 3100 to 3300 cm⁻¹ is assigned to the O–H stretching of hydroxyl and carboxyl groups (Coury and Dillner, 2008). The peak at 1717 cm⁻¹ represents the C=O stretching of aldehydes, ketones, and carboxylic acids. The peaks at 1622 and 1278 cm⁻¹ show good correlation and both are assigned to the -ONO₂ stretching (Liu et al., 2012; Jia and Xu, 2014). The characteristic absorption band at 1500-1350 cm⁻¹ is the C–O stretching and O–H bending of COOH group (Ofner et al., 2011). The absorption peak of sulfate exists in the range of 1200-1000 cm⁻¹ (Wu et al., 2013). The band at 1100 cm⁻¹ in the IR spectra can be attributed to the sulfate group in organic compounds and sulfate. It has been confirmed that the S=O absorption band in organic sulfate monoesters appears around 1040-1070 cm⁻¹ (Chihara, 1958). Although, more studies on bands assignments in organosulfates are not currently available from the literature for further comparison, the 1100 cm⁻¹ band from the current FTIR study can reasonably be assigned to S=O in the sulfate group.

The intensities of most absorption bands, such as O-H at 3100-3300 cm⁻¹, C=O at 1717 cm⁻¹, -ONO₂ at 1622 and 1278 cm⁻¹, and C-H at 2930 cm⁻¹, have similar trends with the change of SOA yield for initial SO₂ concentrations between 11 and 105 ppb. However, the band of sulfate at 1100 cm⁻¹ in IR spectra increases with the rise of initial SO₂ concentration rather than the SOA yield, which suggests the formation of sulfate group in organic compounds and sulfate product from SO₂ photooxidation since, only the relative difference in the intensities of FTIR peaks were studied here. The relative intensity of the band at 1100 cm⁻¹ increased by 1.8 times when the initial SO₂ concentration rose from 0 to 44 ppb, while the band increased by 7.2 times when the initial SO₂ concentration was 105 ppb. The intensity of 1100 cm⁻¹ band grew slowly at low SO₂ concentrations due to the decrease in the formation of aerosols. To clearly show the amount of sulfate group and sulfate in aerosols, the intensity of the band at 1100 cm⁻¹ and the amount of SO₄²⁻ were compared in the same aerosol mass, as shown in Figure 5. The relative intensity was set to 1 when the initial SO₂ concentration was 44.3 ppb.

The relative band intensities at 1100 cm^{-1} , which represented the intensity of both SO_4^{2-} and the sulfate group in organic compounds, increased approximately in a linear form with the increase of initial SO_2 concentration ($R^2=0.91$). If the 1100 cm^{-1} band originated from SO_4^{2-} only, the change of the band intensity would be consistent with SO_4^{2-} concentration in unit mass of aerosols. Figure 5 shows the inconsistency between the FTIR band at 1100 cm^{-1} and the amount SO_4^{2-} as the initial SO_2 concentration, which implies that the 1100 cm^{-1} band originated not only from SO_4^{2-} , but also from other organosulfur compounds. These include organosulfates, which also have the S=O bond, and might contribute to the 1100 cm^{-1} band in the FTIR spectrum. The gap between the FTIR band at 1100 cm^{-1} and SO_4^{2-} can be attributed to the formation of organosulfates.

The composition of the cyclohexene SOA was examined with HR-MS using negative ion mode ESI and the mass spectrum was recorded at a resolution of 10^5 (Figure 6). The OH addition to the C=C bond produces an alkyl peroxy (RO_2) radical that can react with NO to yield organonitrates (Perring et al., 2013). Although the formation of organonitrates was highly expected, there was no evidence of the presence of N-containing compounds from the main peaks of Figure 6, indicating that organonitrates would be formed at very low concentrations, if at all. A similar conclusion could be observed from Figure 4, when noticing that the $-\text{ONO}_2$ stretching peaks at 1622 and 1230 cm^{-1} have very low intensities. The presumed low concentrations of organonitrates might be due to the low concentration of NO when SOA was formed. RO_2 radicals also react with NO_2 to form peroxy nitrates (RO_2NO_2) on time scales comparable to RONO_2 formation. However, RO_2NO_2 are thermally labile and rapidly dissociate at ambient temperatures (Perring et al., 2013). Organosulfates were identified in the particle phase from the chamber experiments. Accurate mass fittings for measured ions of organosulfates in ESI negative ion mode are given in Table 2. As shown in Figure 6 and Table 2, 10 different organosulfates were successfully detected and identified from cyclohexene SOA. These results not only first prove the formation of organosulfates from cyclohexene photooxidation at high- NO_x condition in the presence of SO_2 , but also provide evidence and reference for organosulfates identification by FTIR-IC joint technique. A deprotonated molecular ion at $m/z=195.03322$ ($\text{C}_6\text{H}_{11}\text{O}_3\text{S}^-$) had the maximum content (more than 60%) of all the organosulfates detected in our study. Its intensity was 6.5 times higher than that of the second highest abundant organosulfate. The intermediate product of cyclohexene reaction with OH, i.e., $\text{CH}(\text{O})\text{CH}_2\text{CH}_2\text{CH}_2\text{CH}_2\dot{\text{C}}\text{HOH}$, has a hydroxyl group, and the organosulfate product ($m/z=195.03322$) would likely form from the intermediate product, not from end product.

The mass spectra show a great abundance of peaks, detected as deprotonated molecular ions ($\text{M}-\text{H}^-$) formed via proton abstraction. Most cyclohexene SOA contained carboxylic acid and/or aldehyde moieties. The products of the reaction of OH radicals with cyclohexene in the presence of NO were investigated and were identified as cyclic 1,2-hydroxynitrates and 1,6-hexanedial (Aschmann et al., 2012). These products could not be detected by ESI-HR-MS in our study. Aldehydes could be oxidized by OH radicals to form carboxyls, which have been intensively identified in previous studies (Cameron et al., 2002; Goldsmith et al., 2012). The 1,6-hexanedial might be further oxidized in the atmospheric photooxidation reactions to form the 1,6-adipic acid ($\text{C}_6\text{H}_{10}\text{O}_4$) and 6-oxohexanoic acid ($\text{C}_6\text{H}_{10}\text{O}_3$), which were both observed in our MS spectra. In addition to the C_6 compounds observed in this study, a $\text{C}_5\text{H}_7\text{O}_3^-$ ion was detected with higher abundance than the C_6 compounds. Although the formation of this C_5 compound might be due to a carbonyl cleavage from a six-carbon atoms chain, a proper

mechanism for its formation could not be determined. A C₄ compound was also detected in this study, likely as a result of a carbonyl cleavage from a C₅ compound. However, there was no evidence of the formation of compounds with less than four carbon atoms.

The ESI-HR-MS spectra of particles formed from different initial SO₂ concentrations are shown in Figure S6. We found no obvious difference in the composition and response of organosulfates with different initial SO₂ concentrations. The relative intensity of m/z = 97, which corresponds to sulfate, was set to 100% in both ESI-HR-MS spectra. The relative intensities of the organosulfates peaks in both spectra were almost unchanged regardless of the initial SO₂ concentration. Our result is consistent with the results of Minerath et al. and Hatch et al. who observed an increase in organosulfates yield with increasing sulfate concentration (Minerath and Elrod, 2009; Hatch et al., 2011). These observations demonstrate that particle sulfate content is likely a key parameter influencing organosulfates formation. **Given the consistency between sulfates and organosulfates concentrations as shown in Figure S6 and S7, it is most likely that as the amount of sulfate increase with increasing initial SO₂ concentration, the concentration of organosulfates will also increase. Comparing HR-MS data when initial concentrations were 0 ppb and 236 ppb reveals that the bands representing organosulfates do not appear at 0 ppb of SO₂. Peaks at m/z larger than 150 were undetectable at initial SO₂ concentration of 0 ppb, while products without sulfur peaked at both concentrations, with the only difference being their intensities. This implies that the process of SOA formation strongly depends on initial SO₂ concentrations.**

4 Conclusion

We report a series of chamber experiments studies on the formation of secondary aerosols from the mixture of cyclohexene and SO₂. The experiments were based on Fourier transform infrared spectrometer, ion chromatography and electrospray ionization high-resolution quadrupole mass spectrometer, and were performed under NO_x conditions. Although new particle formation was found to be enhanced with increasing SO₂ concentration, the yield of SOA was not enhanced for all SO₂ concentrations between 0 and 105 ppb. The SOA formation decreased at first and then was enhanced for all SO₂ concentration above 40 ppb. Both acid-catalysis and competitive OH reactions with cyclohexene and SO₂ were found to have important effects on the SOA formation and hence, should be taken into account in SOA simulation models or air quality models for a better understanding of haze pollution. The formation of organosulfates, an important part of atmospheric organic aerosol components, was first observed from cyclohexene SOA. The formation of organosulfates has a great significance for the particulate matter formation under high SO₂ concentrations in the atmosphere.

Acknowledgments

This work was supported by National Natural Science Foundation of China (91644214, 21577080, 41375129), Shenzhen Science and Technology Research and Development Funds, China (JCYJ20150402105524052), and the “Strategic Priority Research Program (B)” of the Chinese Academy of Sciences (XDB05010104).

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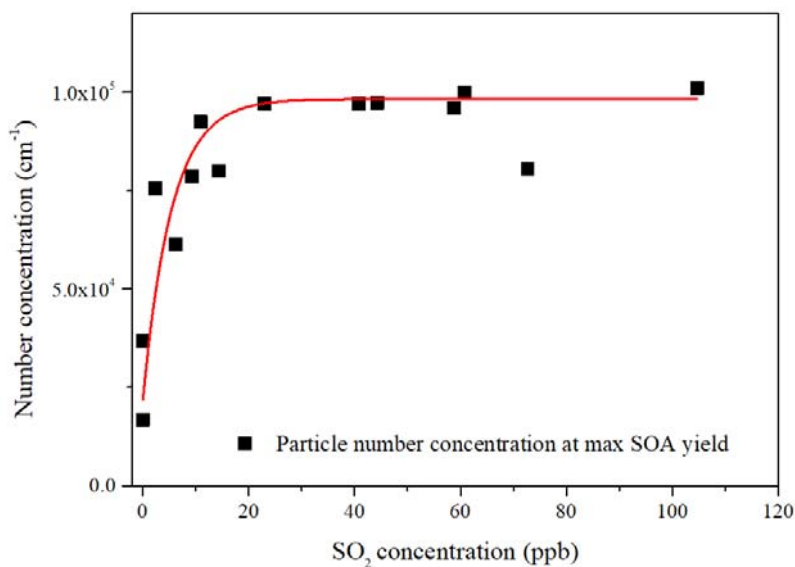
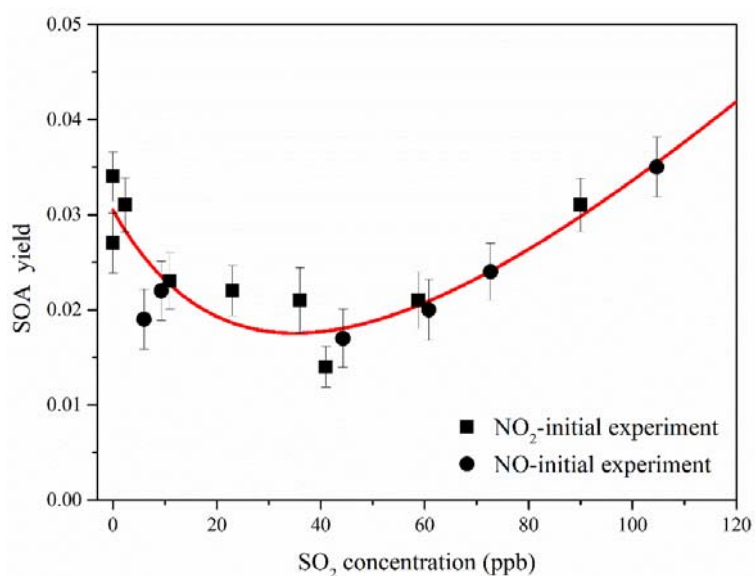


Figure 1: Particle number concentrations of SOA in the photooxidation of the cyclohexene/NO_x/SO₂ system with different initial SO₂ concentrations.



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Figure 2: SOA yields of cyclohexene photooxidation in the presence of NO_x at different initial SO₂ concentrations. Solid line is the least-square fitting to the data. The error bars were determined on the basis of propagation of uncertainties arising in the ΔHC measurements, including GC calibration uncertainties propagation and the variance in the initial cyclohexene measurements.

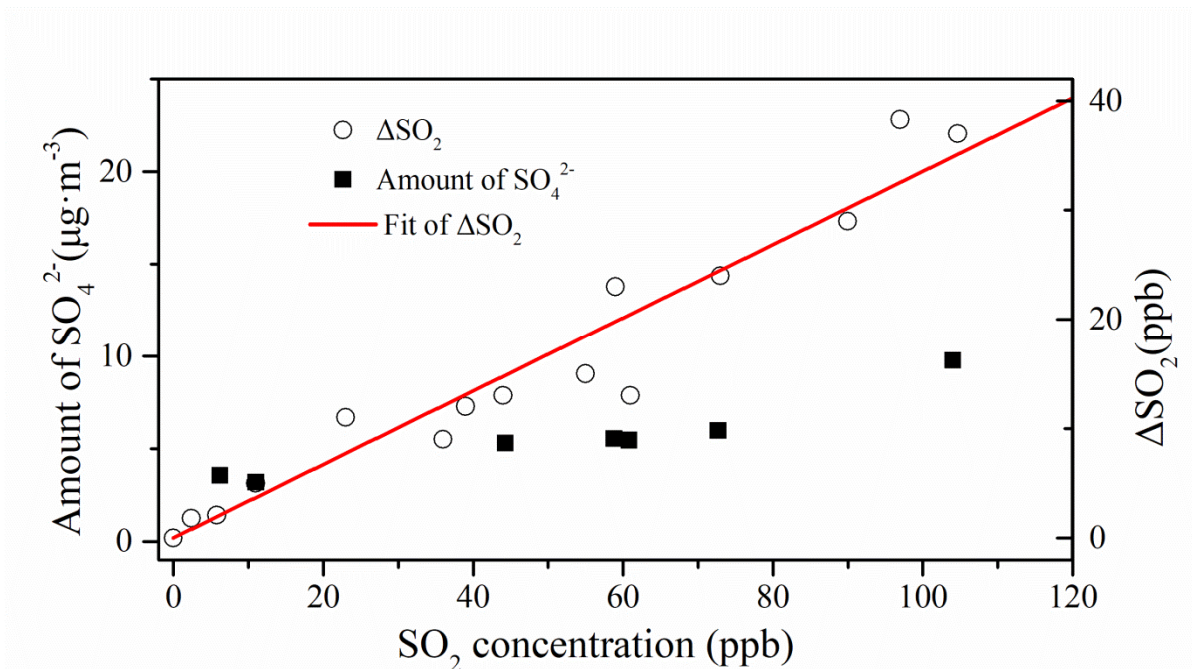


Figure 3: The amount of SO_4^{2-} in particle phase and the consumption of SO_2 (ΔSO_2) with different initial SO_2 concentrations.

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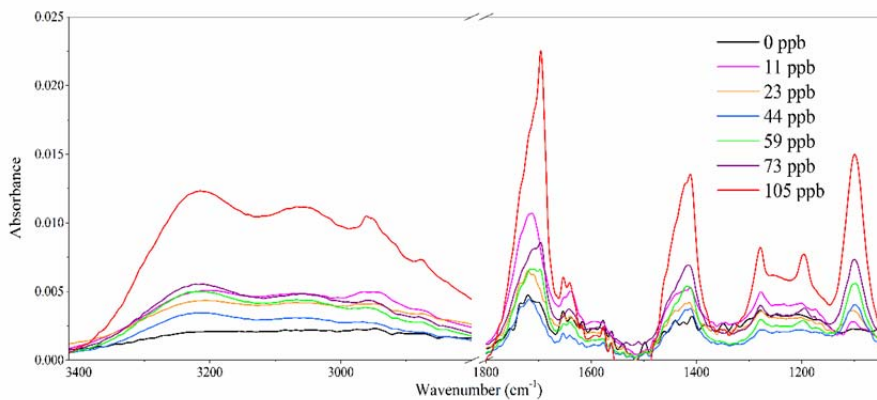


Figure 4: IR spectra of aerosols from the cyclohexene/NOx/SO₂ system under different SO₂ concentrations.

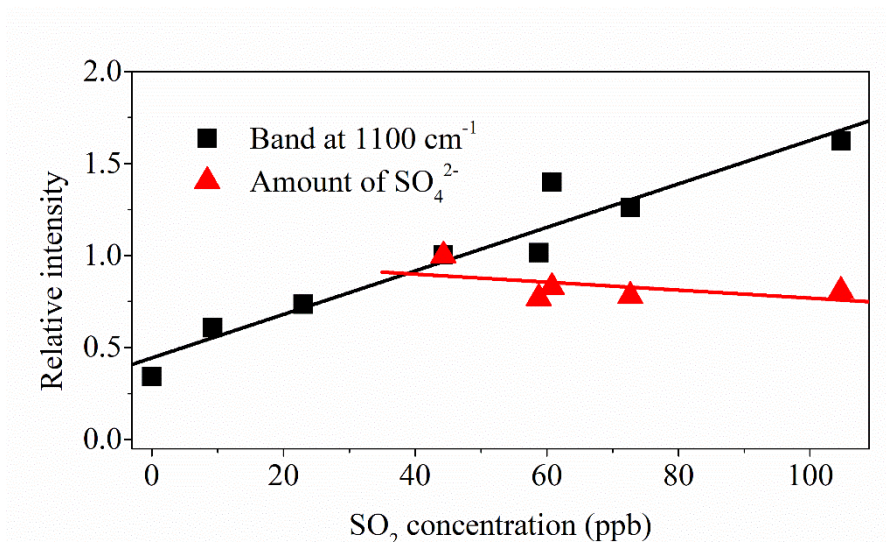


Figure 5: The relative intensity of the FTIR band at 1100 cm⁻¹ (square) and the amount of SO₄²⁻ (triangle) normalized to SOA mass. The 1100 cm⁻¹ band intensity and the amount of SO₄²⁻ were divided by the formed SOA mass, firstly. Subsequently, the results of both FTIR band at 1100 cm⁻¹ and the amount of SO₄²⁻ divided by SOA mass were set to 1 at the experiment which initial SO₂ concentration was 44.3 ppb.

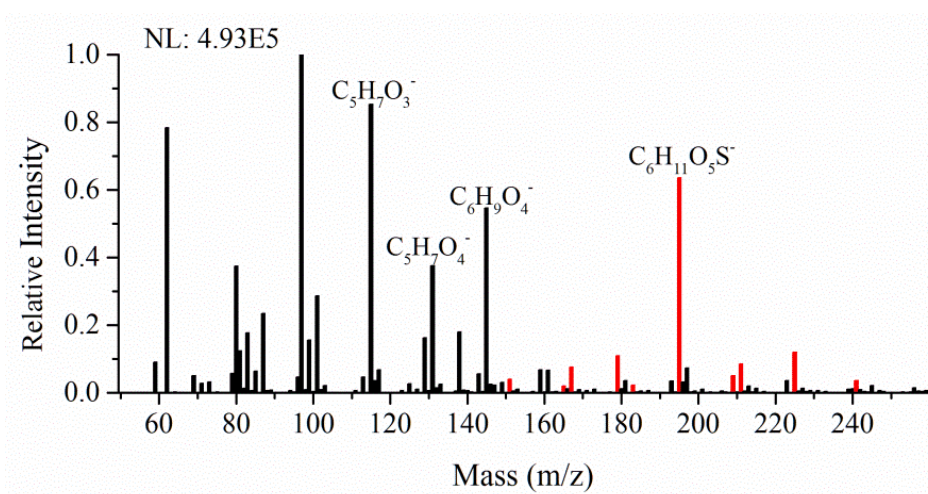


Figure 6: Negative ion mode ESI mass spectrum of SOA generated from the photooxidation of cyclohexene in the presence of SO₂. Red peaks correspond to organic compounds containing the sulfate group. The mass resolution is 10⁵.

Table 1 Experimental conditions for the photooxidation of cyclohexene/NO_x/SO₂ system. All experiments were performed under dry conditions (relative humidity < 10 %). ΔM_0 is the produced organic aerosol mass concentration and Y is the SOA yield.

Exp.	T (K)	SO ₂ (ppb)	cyclohexene (ppb)	NO _x (ppb)	cyclohexene/N Ox	ΔM_0 ($\mu\text{g m}^{-3}$)	Y (%)
1 ^b	308	0.0	596	122.0	4.9	57.0	2.66
2 ^b	305	0.0	651	93.7	6.9	79.7	3.40
3 ^b	309	2.4	553	95.7	5.8	62.6	3.15
4 ^a	307	5.8	612	92.7	6.6	41.0	1.87
5 ^a	309	9.3	599	93.5	6.4	48.1	2.23
6 ^b	309	11.0	574	94.7	6.1	47.1	2.28
7 ^b	309	23.0	514	90.5	5.7	42.6	2.30
8 ^b	305	36.6	665	99.7	6.7	96.3	2.01
9 ^b	308	40.8	472	91.4	5.2	22.6	1.33
10 ^a	308	44.3	592	98.6	6.0	35.3	1.66
11 ^b	305	55.0	497	113.0	4.4	77.3	2.16
12 ^b	308	58.8	577	96.7	6.0	44.3	2.13
13 ^a	309	60.8	626	102.0	6.1	43.9	1.95
14 ^a	308	72.7	581	98.4	5.9	49.2	2.35
15 ^b	306	90.0	543	99.6	5.4	102.0	2.62
16 ^a	309	104.7	608	93.7	6.5	77.1	3.52
17 ^{bc}	305	236.0	1048	198.0	5.3	-	-
18 ^{bc}	306	93.7	1235	215	5.7	-	-

5 ^a: the experiment was initiated by NO.

^b: the experiment was initiated by NO₂.

^c: the formed particles were detected by ESI-HR-MS.

Table 2. Accurate mass fittings for main products and measured organosulfates ions in ESI negative ion mode from cyclohexene photooxidation in the presence of SO₂ under high-NO_x conditions

Measured ^a m/z	Ion	Proposed Ion Formula	Delta ^b (ppm)	RDB ^c
115.03942		C ₃ H ₇ O ₃ ⁻	-5.628	2
145.05019		C ₆ H ₉ O ₄ ⁻	-3.048	2
131.03444		C ₅ H ₇ O ₄ ⁻	-4.136	2
101.06006	(M-H) ⁻	C ₅ H ₉ O ₂ ⁻	-7.351	1
87.04433		C ₄ H ₇ O ₂ ⁻	-9.453	1
129.05515		C ₆ H ₉ O ₃ ⁻	-4.397	2
99.04439		C ₃ H ₇ O ₂ ⁻	-7.702	2
Organosulfates				
195.03322		C ₆ H ₁₁ O ₅ S ⁻	-0.243	1
225.00771		C ₆ H ₉ O ₇ S ⁻	1.171	2
179.00181		C ₅ H ₇ O ₅ S ⁻	-0.0879	2
211.02828		C ₆ H ₁₁ O ₆ S ⁻	0.464	1
167.00167	(M-H) ⁻	C ₄ H ₇ O ₅ S ⁻	-1.780	1
209.01257		C ₆ H ₉ O ₆ S ⁻	0.182	2
151.00658		C ₄ H ₇ O ₄ S ⁻	-3.130	1
241.00278		C ₆ H ₉ O ₈ S ⁻	1.738	2
182.99667		C ₄ H ₇ O ₆ S ⁻	-1.158	1
164.98594		C ₄ H ₅ O ₅ S ⁻	-2.287	2

^a Sort by abundance intensity.

^b Delta: label the peak with the difference between the theoretical and measured m/z.

5 ^c RDB: ring and double bond equivalent.