

1 **Responses to Referee # 1**

2 We thank the reviewer for the careful review of our manuscript. The comments and suggestions are
3 greatly appreciated. All the comments have been addressed. In the following, please find our
4 responses to the comments one by one and the corresponding revisions made to the manuscript. The
5 original comments are shown in italics. The revised parts of the manuscript are highlighted.

6 **Anonymous Referee #1**

7 Received and published: 26 April 2017

8 *This manuscript describes the competing roles of NO_x and SO₂ on SOA formation of α -pinene and*
9 *limonene. The ability of SO₂ to enhance seed aerosol surface area appears to be a dominant factor,*
10 *and that enhancing seed aerosol reduces the NO_x suppression of SOA yields, at least in some*
11 *monoterpenes. The authors use their AMS data to determine the role of organic nitrates in SOA, and*
12 *find that organic nitrates account for a substantial fraction of the SOA mass. Overall, this is an*
13 *interesting piece of work, and warrants publication in the ACP following some revision.*

14 *Major Comments*

15 *The nature of the experimental design was not so much to look at the impact of SO₂ – but to look at*
16 *the role of a sulfate seed aerosol. From the manuscript, my interpretation is that the SO₂ additions*
17 *were used to nucleate (inorganic) seed aerosol. Was there any SO₂ left over to impact VOC oxidation?*
18 *It is not clear to me if the SO₂ additions really paralleled the NO_x additions, because the*
19 *experimental design was different. That’s not to say that these aren’t valuable experiments that add to*
20 *the literature! I merely question whether this was truly an ‘SO₂ addition’ rather than a ‘sulfate*
21 *aerosol addition’ to VOC oxidation experiments.*

22 **Response:**

23 We thank the reviewer for the supporting remarks.

24 In the experiments with SO₂ added, SO₂ concentration decreased slowly and most of SO₂ was still left
25 (typically around 8 ppb) at the end of an experiment because of the low reactivity of SO₂ with OH
26 ($\sim 2 \times 10^{-12}$ molecules⁻¹ cm³ s⁻¹ at 298 K). SO₂ time series in a typical experiment are shown in a newly
27 added figure (Fig. S2). Therefore, the experiments with SO₂ not only included the effect of sulfate
28 formed from SO₂ oxidation as seed but also the potential role of SO₂ on VOC oxidation, although the
29 role on VOC oxidation turned out to be likely not significant. We have added the follow sentence in
30 the revised manuscript to clarify this point.

31 “SO₂ concentration decayed slowly in the experiments with SO₂ added and most of the SO₂ was still
32 left (typically around 8 ppb from initial 15 ppb) at the end of an experiment due to its low reactivity
33 with OH. Typical SO₂ time series in high SO₂ experiments are shown in Fig S2.”

34 Although the SO₂ addition did not exactly parallel NO_x addition, by adding SO₂ and inducing
35 nucleation first, we can make sure that in high SO₂ conditions enough nucleated particles were
36 represent for the oxidation products to condense on once VOC oxidation started. Otherwise, it would
37 be unclear whether the low SOA yield at high NO_x was due to missing nucleation or lack of
38 condensable products. In this way, we can somewhat differentiate the role of promoting nucleation
39 from the role of affecting the condensable products from VOC oxidation.

40 In the revised manuscript, we have added the following sentence.

41 “Adding SO₂ first and initializing nucleation by SO₂ photooxidation ensured that enough nucleating
42 particles were present when VOC oxidation started.”

43 The effect of SO₂ on VOC oxidation is provided in the response to one similar comment below (Pg. 7,
44 lines 205-209).

45 *SOA yield is influenced strongly by OA mass. The authors plot SOA yield versus OH dose, which is*
46 *certainly a useful figure to see – but it is hard to compare the SOA yields if the SOA mass has not*
47 *been accounted for. The authors need to also show SOA yield versus OA mass so that the readers can*
48 *contrast the relationships to other studies. It would be useful to compare the SOA yields to other*
49 *studies: how do the yield values compare to other measurements of OH oxidation of a-pinene? This*
50 *will allow readers to place the studies in context.*

51 **Response:**

52 We have accepted the reviewer’s suggestion. In the revised manuscript, we have added a figure of
53 SOA yield versus OA mass concentration (Fig. S8 and S9) and compared the SOA yield in this study
54 to previous studies. We have also discussed other factors influencing SOA yield.

55 *The results of SO₂ and NO_x effects on SOA yield are consistent with the Sarrafzadeh and Eddingsaas*
56 *studies, which found that the presence of seed aerosol suppresses the ‘NO_x effect’ on SOA yield.*
57 *However, they contradict previous studies (e.g. Ng et al. 2007, Presto et al. 2005). The authors need*
58 *to do a better job of contrasting their studies – they attribute the difference to a vague collection of*
59 *parameters (e.g. NO:NO₂ ratio, OH concentrations, etc.). It would be extremely helpful if the authors*
60 *could synthesize the information (i.e. put numbers on those parameters) to help readers understand*
61 *the differences in experimental conditions across the studies. A table would be particularly helpful.*

62 **Response:**

63 We have accepted the reviewer’s suggestions. In the revised manuscript, we have added one table
64 summarizing the reaction conditions of previous studies (Table. S2) and elaborated the discussion
65 related to the difference between our study and previous studies.

66 “...The reaction conditions of this study often differ from those described in the literature (see
67 Table S2).

68 The difference in these conditions can result in both different apparent dependence on specific
69 parameters and the varied SOA yield. For example, SOA yield from α -pinene photooxidation at low
70 NO_x in this study appeared to be much lower than that in Eddingsaas et al. (2012a). The difference
71 between the SOA yield in this study and some of previous studies and between the values in the
72 literature can be attributed to several reasons: 1) RO_2 fates may be different. For example, in our study
73 at low NO_x , RO_2+NO account for a large fraction of RO_2 loss while in Eddingsaas et al. (2012a)
74 RO_2+HO_2 is the dominant pathway of RO_2 loss. This difference in RO_2 fates may affect oxidation
75 products distribution. 2) The organic aerosol loading of this study is much lower than that some of
76 previous studies (e.g., Eddingsaas et al. (2012a)) (see Fig. S9). SOA yields in this study were also
77 plotted versus organic aerosol loading to better compare with previous studies (Fig. S8 and S9). 3)
78 The total particle surface area in this study may also differ from previous studies, which may
79 influence the apparent SOA yield due to vapor wall loss (the total particle surface area is often not
80 reported in many previous studies to compare with). 4) RH of this study is different from many
81 previous studies, which often used very low RH (<10%). It is important to emphasize that reaction
82 conditions including the NO_x as well as SO_2 concentration range and RH in this study were chosen to
83 be relevant to the anthropogenic-biogenic interactions in the ambient atmosphere. In addition,
84 difference in the organic aerosol density used in yield calculation should be taken into account. In this
85 study, SOA yield was derived using a density of 1 g cm^{-3} to better compare with many previous
86 studies (e.g., (Henry et al., 2012)), while in some other studies SOA yield was derived using different
87 density (e.g., 1.32 g cm^{-3} in Eddingsaas et al. (2012a)).”

88 *Lines 118: the use of the HR-ToF-AMS to derive elemental ratios uses the older Aiken method.*
89 *However, as the authors note, the newer 2015 approach corrects some underestimation. Because*
90 *readers may wish to compare results across studies in the future, it is appropriate and prudent to*
91 *update the results to the newer calculations.*

92 **Response:**

93 We have calculated the H/C and O/C data using the newer approach by Canagaratna et al. (2015) and
94 compared them with the data derived from the older method (Aiken et al., 2007) (Fig. S1). The H/C
95 derived using the newer method strongly correlated with that derived using older method and just
96 increased by 27%. Similarly, O/C just increased by 11%. In the revised manuscript, we have
97 discussed this difference.

98 “The H/C and O/C were also derived using the newer approach by Canagaratna et al. (2015) and
99 compared with the data derived from the Aiken et al. (2007) method. The H/C values derived using
100 the Canagaratna et al. (2015) method strongly correlated with the values derived using Aiken et al.

101 (2007) method (Fig. S1) and just increased by 27% as suggested by Canagaratna et al. (2015). Similar
102 results were found for O/C and there was just a difference of 11% in O/C. Since only relative
103 difference in elemental composition of SOA is studied here, only the data derived using Aiken et al.
104 (2007) method are shown as the conclusion was not affected by the methods chosen.”

105 *Line 128: the authors note that they account for particle wall loss and dilution loss, but not for vapor*
106 *wall loss. Recent papers have shown this to be a chemically-dependent and substantial effect on SOA*
107 *yields, and most rigorous SOA yield work now accounts for these effects. How will ignoring vapor*
108 *wall loss influence the results – and the interpretation thereof?*

109 **Response:**

110 The wall loss of vapors causes an under-estimate of the SOA yield. In the revised manuscript, we
111 have estimated the influence of the vapor wall loss on SOA yield using the measured wall loss rate of
112 vapors. And we have added a section to address the influence of vapor wall loss.

113 **“Wall loss of organic vapors**

114 The loss of organic vapors on chamber walls can influence SOA yield (Kroll et al., 2007;
115 Zhang et al., 2014; Ehn et al., 2014; Sarrafzadeh et al., 2016; McVay et al., 2016; Nah et al., 2016;
116 Matsunaga and Ziemann, 2010; Ye et al., 2016; Loza et al., 2010). The wall loss rate of organic
117 vapors in our chamber was estimated by following the decay of organic vapor concentrations after
118 photooxidation was stopped in the experiments with low particle surface area ($\sim 5 \times 10^{-8} \text{ cm}^2 \text{ cm}^{-3}$) and
119 thus low condensational sink on particles. Such method is similar to the method used in previous
120 studies (Ehn et al., 2014; Sarrafzadeh et al., 2016; Krechmer et al., 2016; Zhang et al., 2015). A high-
121 resolution time-of-flight chemical ionization mass spectrometer (HR-ToF-CIMS, Aerodyne Research
122 Inc.) with nitrate ion source ($^{15}\text{NO}_3^-$) was used to measure semi/low-volatile organic vapors. The
123 details of the instrument were described in our previous studies (Ehn et al., 2014; Sarrafzadeh et al.,
124 2016). The decay of vapors started from the time when the roof of the chamber was closed. The data
125 were acquired at a time resolution of 4 s. A typical decay of low-volatile organics is shown in Fig. S3
126 and the first-order wall loss rate was determined to be around $6 \times 10^{-4} \text{ s}^{-1}$.

127 The SOA yield was not directly corrected for the vapor wall loss, but the influence of vapor
128 wall loss on SOA yield was estimated using the method in the study of Sarrafzadeh et al. (2016) and
129 the details of the method are described therein. Briefly, particle surface and chamber walls competed
130 for the vapor loss (condensation) and the condensation on particles led to particle growth. The fraction
131 of organic vapor loss to particles in the sum of the vapor loss to chamber walls and to particles (F_p)
132 was calculated. The vapor loss to chamber walls was derived using the wall loss rate. The vapor loss
133 to particles was derived using particle surface area concentration, molecular velocity and an
134 accommodation coefficient α_p (Sarrafzadeh et al., 2016). $1/F_p (f_{\text{corr}})$ provides the correction factor to

135 obtain the “real” SOA yield. f_{corr} is a function of particle surface area concentration and
136 accommodation coefficient as shown in Fig. S4. Here a range of 0.1-1 for α_p was used, which is
137 generally in line with the ranges of α_p found by Nah et al. (2016) by fitting a vapor-particle dynamic
138 model to experimental data. At a given α_p , the higher particle surface area, the lower f_{corr} and the
139 lower the influence of vapor wall loss are because most vapors condense on particle surface and vice
140 versa. At a given particle surface area, f_{corr} decreases with α_p because at higher α_p a larger fraction of
141 vapors condenses on particles. An average molecular weight of 200 g/mol was used to estimate the
142 influence of vapor wall loss. For the aerosol surface area range in most of the experiments in this
143 study, the influence of vapor wall loss on SOA yield was relatively small (<~40% for particle surface
144 area larger than $3 \times 10^{-6} \text{ cm}^2 \text{ cm}^{-3}$, Fig. S4). Yet, for the experiments at high NO_x and low SO_2 for α -
145 pinene and limonene, the influence of vapor wall loss on SOA can be high due to the low particle
146 surface area, especially at lower α_p .”

147 *In the Introduction, the authors do a good job of summarizing the reasons why such a study would be*
148 *interesting. Much of the discussion focuses on the role of NO_x on SOA yields – this is reasonable as*
149 *most of the literature has focused on that problem! However, there is some relatively recent literature*
150 *regarding the role of SO_2 in affecting SOA chemistry and monoterpene OH oxidation that the authors*
151 *should consider. In particular:*

152 *Photooxidation of cyclohexene in the presence of SO_2 : SOA yield and chemical composition.*

153 *Shijie Liu, Long Jia, Yongfu Xu, Narcisse T. Tsona, Shuangshuang Ge, and Lin Du. Atmos. Chem.*
154 *Phys. Discuss., doi:10.5194/acp-2017-30, 2017*

155 *Synergetic formation of secondary inorganic and organic aerosol: effect of SO_2 and NH_3 on particle*
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158 *2016*

159 *Formation of secondary aerosols from gasoline vehicle exhaust when mixing with SO_2 . T. Liu, X.*
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163 *Anthropogenic Sulfur Perturbations on Biogenic Oxidation: SO_2 Additions Impact Gas- Phase OH*
164 *Oxidation Products of α - and β -Pinene. Beth Friedman, Patrick Brophy, William H. Brune, and*
165 *Delphine K. Farmer. Environmental Science & Technology 2016 50 (3), 1269-1279. DOI:*
166 *10.1021/acs.est.5b05010*

167 *Is there any evidence for organic sulfates in the SOA from the AMS data? This has been a subject of*
168 *some debate in the literature, and an additional datapoint would be useful. This may also clarify the*
169 *role of acid catalysis, as I believe that has been linked to the formation of organic sulfates.*

170 **Response:**

171 We thank the reviewer for raising these papers. In the revised manuscript, we have enriched the
172 discussion on the role of SO₂ by including some of the papers.

173 From our AMS data, we did not find evidence of organic sulfate. For SOA formed at high SO₂, we
174 found no significant organic fragments containing sulfur. Also the fragment CH₃SO₂⁺ from organic
175 sulfate suggested by Farmer et al. (2010) was not detectable in our data. We found that the pattern of
176 sulfate in mass spectra had no significant difference from the pattern of pure ammonium sulfate.
177 However, we would like to note that AMS has very limited capability to differentiate organic sulfate
178 and inorganic sulfate (Farmer et al., 2010).

179 Moreover, according to the literature, organic sulfate is mainly formed by aqueous reaction of sulfate
180 with organics. In the conditions of our study, there was no aqueous phase as we stated based on the
181 AIM model. Therefore, experimental conditions in our study did not favor the formation organic
182 sulfate.

183 In the revised manuscript, we have clarified these results.

184 “In addition, from the AMS data of SOA formed at high SO₂ no significant organic fragments
185 containing sulfur were found. Also the fragment CH₃SO₂⁺ from organic sulfate suggested by Farmer
186 et al. (2010) was not detected in our data. The absence of organic sulfate tracers is likely due to the
187 lack of aqueous phase in aerosol particles in this study. Therefore, the influence of SO₂ on gas phase
188 chemistry of organics and further on SOA yield via affecting gas phase chemistry is not important in
189 this study.”

190 Minor Comments

191 *Line 136. The authors note an average RH of 28-42% for the experiments. This seems like a relatively*
192 *large range: will this affect the SOA yields, or interpretation of the data?*

193 **Response:**

194 The average RH was in the range of 28-42% taking into account all experiments. Actually, except one
195 experiment, the average RH was in the range of 28-34%. For particle phase reactions, the particle
196 water content absorbed by organic aerosol in the range of 28-42% RH is low and the difference of
197 water content between 28% and 42 % is very minor (typically <~2% of the particle volume based on
198 our hygroscopic growth measurement). The RH variations are not expected to significantly change the
199 particle phase chemistry. Moreover, since water vapor is abundant and in excess in the gas phase, the

200 RH variations are not expected to significant change gas phase chemistry either. Therefore, we do not
201 expect that would significantly change the SOA yield. In the revised manuscript, we have clarified
202 this point.

203 “The average RH for the period of monoterpene photooxidation was 28-34% except for one
204 experiment with average RH of 42% RH.”

205 *Re: Discussion of SO₂ effects. The authors dominantly attribute the enhancement of SOA by SO₂ to*
206 *increased particle surface area, or perhaps to acid catalysis. These seem like extremely likely reasons;*
207 *however, there is one study that suggests that SO₂ will influence gas-phase oxidation products*
208 *(Friedman et al.), which could also be a confounding factor unless all of the SO₂ is in the particle*
209 *phase before VOC oxidation commences... This would be a useful clarification.*

210 **Response:**

211 The influence of SO₂ on gas phase oxidation is likely to be trivial in this study for two reasons. Firstly,
212 the reactivity of SO₂ with OH is very low (2×10^{-12} vs. 5.3×10^{-11} molecules⁻¹ cm³ s⁻¹ for α -pinene with
213 OH) and SO₂ only accounts for a very small fraction of the OH loss (typically ~2% in the beginning of
214 an experiment). Secondly, the OH concentration is 2-3 orders of magnitude lower than those in the
215 PAM chamber used by Friedman et al. (2016). Therefore, either the change in OH/HO₂ ratio or SO₃
216 concentration, which is attributed to the reason of changed oxidation products by Friedman et al.
217 (2016), is much lower in our experiments. In the revised manuscript, we have added a brief discussion
218 of the effect of SO₂ on gas phase oxidation of monoterpenes in this study as follows.

219 “SO₂ has been proposed to also affect gas phase chemistry of organics by changing the HO₂/OH or
220 forming SO₃ (Friedman et al., 2016). In this study, the effect of SO₂ on gas phase chemistry of
221 organics was not significant because of the much lower reactivity of SO₂ with OH compared with α -
222 pinene and limonene (Atkinson et al., 2004, 2006; Atkinson and Arey, 2003) and the low OH
223 concentrations (2-3 orders of magnitude lower than those in the study by Friedman et al. (2016)).
224 Moreover, reactions of RO₂ with SO₂ was also not important because the reaction rate constant is very
225 low ($< 10^{-14}$ molecule⁻¹ cm³ s⁻¹) (Lightfoot et al., 1992; Berndt et al., 2015). In addition, from the AMS
226 data of SOA formed at high SO₂ no significant organic fragments containing sulfur were found. Also
227 the fragment CH₃SO₂⁺ from organic sulfate suggested by Farmer et al. (2010) was not detected in our
228 data. The absence of organic sulfate tracers is likely due to the lack of aqueous phase in aerosol
229 particles in this study. Therefore, the influence of SO₂ on gas phase chemistry of organics and further
230 on SOA yield via affecting gas phase chemistry is not important in this study.”

231 *Technical comments.*

232 *Line 26: should read “compared to low NO_x”*

233 **Response:** Corrected.

234 *Line 29: should read “SO₂ can compensate for such effects”*

235 **Response:** Corrected.

236 *Introduction: line 34: sentence has repetitive ‘important’: consider removing at least one (e.g. “SOA*
237 *is an important class of atmospheric aerosol” seems like an unnecessary statement for the journal’s*
238 *audience). This adjective is used heavily throughout the introduction (lines 45, 49), and I recommend*
239 *removing or replacing the adjective to improve readability.*

240 **Response:** We have accepted the reviewer’s suggestion. In the revised manuscript, we have removed
241 the “as an important class of atmospheric aerosol”, and removed or replaced “important” where it is
242 necessary.

243 *Line 56: hydroperoxides should be plural*

244 **Response:** Corrected.

245 *Line 57: need comma between ‘NO’ and ‘forming’*

246 **Response:** Corrected.

247 *Line 87: should read “might have either counteracting or synergistic effects on SOA: : :”*

248 **Response:** Corrected.

249 *Line 126: remove the with following ‘multiplied by’*

250 **Response:** Done.

251 *Line 135: should read ‘there was no aqueous..’*

252 **Response:** Corrected.

253 *Line 221, remove comma between ‘that’ and ‘high’*

254 **Response:** Done.

255 *Line 360: should read ‘in the ambient atmosphere’*

256 **Response:** Corrected.

257 **References**

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1 **Responses to Referee # 2**

2 We thank the reviewer for carefully reviewing our manuscript; the comments and suggestions are
3 very helpful and greatly appreciated. All the comments have been addressed. We believe that
4 revisions based on these comment have substantially improved our manuscript. In the following
5 please find our responses to the comments one by one and the corresponding changes made to the
6 manuscript. The original comments are shown in italics. The revised parts of the manuscript are
7 highlighted.

8 Before we start our responses to the reviewer, we would like to briefly clarify the motivation of this
9 manuscript. Our primary goal of this study is to investigate how and to what extent small ambient
10 inorganic trace gases, here NO_x and SO_2 , affect the SOA formation in the ambient
11 anthropogenic-biogenic interactions. We aimed to study the more complex situation found in the
12 ambient atmosphere instead of pure VOC reaction systems. Our main goal is not so much to provide a
13 single absolute value of SOA yield to modelers nor to simply “improve” values of the SOA yields
14 from previous studies because all chamber studies have operational limitations compared to the
15 atmosphere (e.g. wall effects). Often different studies have distinct physical and chemical regimes due
16 to the different operational limitations and diverse experimental conditions e.g. chamber size, radical
17 generation, and photolysis rates. We as well as the community are in the course of addressing the
18 influence of wall loss of vapors. We believe that, rather denying the findings from previous studies,
19 our study provided additional knowledge and insights to existing understanding of SOA formation in
20 certain conditions of the real atmosphere. This is based on the fact that our experiments were
21 conducted under conditions relatively close to the ambient anthropogenic-biogenic interactions,
22 including ambient RH, concentrations of SO_2 , NO_x , and VOC close to ambient levels, natural
23 sunlight, and low surface-to-volume ratio of our large chamber.

24 ***Anonymous Referee #2***

25 *This chamber study investigated the effects of SO_2 and NO_x (NO) on SOA formation from*
26 *photooxidation of α -pinene and limonene. It was found that SO_2 enhanced SOA yield while NO_x*
27 *suppressed SOA yield. The suppression effect of NO_x was attributed to the suppressed new particle*
28 *formation and thus a lack of particle surface area for organics to condense on. The authors concluded*
29 *that SO_2 oxidation produced high number of particles and compensated for the suppression of SOA*
30 *yield by NO_x . SOA composition measured by AMS was also presented and discussed.*

31 *This is an interesting study. The gas- and particle-phase measurements are comprehensive and*
32 *include several important species that have not been typically characterized in previous studies (e.g.,*
33 *OH , HO_2 and RO_2). The experiments appeared to be carefully conducted. However, I have major*
34 *concerns regarding data interpretation and some conclusions in the manuscript.*

35 *One of the central themes of the manuscript is that the suppression effect of NO_x on SOA formation*

36 can be compensated by the presence of SO₂. This conclusion is not accurate based on all the data
37 presented in this manuscript. For α -pinene, it appears that under high SO₂ conditions, the SOA yields
38 under low vs. high NO_x conditions are comparable. However, this is not the case for limonene, where
39 there is still a large difference in SOA yields between low vs. high NO_x conditions in the presence of
40 high SO₂. The manuscript needs to be thoroughly revised to accurately reflect what the data are
41 showing. If one set of data is showing one thing and another set of data is showing the opposite, the
42 authors need to discuss both datasets equally and cannot conclude that SO₂ effect can compensate
43 NO_x effect.

44 **Response:**

45 A key result in our study is that for both α -pinene and limonene, the difference in SOA yield between
46 high NO_x and low NO_x was much reduced in the presence of SO₂, although for limonene the SOA
47 yield at high NO_x and high SO₂ was still lower than the yield at low NO_x. This result indicates that the
48 suppression of SOA yield by NO_x was compensated to a large extent by SO₂. This conclusion holds
49 regardless the difference in the detailed results between α -pinene and limonene system.

50 In the revised manuscript, we have revised the conclusion to better represent the results from both
51 α -pinene and limonene cases. The following sentence has been revised in the conclusion part:

52 “SO₂ oxidation produced high number concentration of particles and compensated for the suppression
53 of SOA yield by NO_x to a large extent.”

54 The abstract has been revised accordingly as follows.

55 “However, in the presence of SO₂ which induced high number concentration of particles after
56 oxidation to H₂SO₄, the suppression of the mass yield of SOA by NO_x was completely or partly
57 compensated.”

58 *The authors concluded that the suppression effect of NO_x on SOA yields is mainly due to suppression
59 of nucleation (absence of particle surface area as condensation sink) rather than decrease of
60 condensable materials. If particle surface area plays a role, this will point to the importance of loss
61 process of oxidation products via chemical reactions and/or chamber wall loss. However, the effect of
62 loss of organic vapors on chamber walls is not considered in this study. Nevertheless, previous studies
63 on α -pinene oxidation suggested that SOA yield is independent of particle surface area. In this regard,
64 the interpretation that the suppression effect of NO_x arises from a lack of particle surface area
65 appears to be at odds with previous studies. All in all, it is not clear how the absence of particle
66 surface area can explain the suppressed SOA yields under high NO_x condition in this study.*

67 **Response:**

68 In general, condensable SOA materials are chemically produced in the gas-phase and nucleation and
69 condensation on particle surface constitute sinks for the SOA materials besides other sinks e.g. wall
70 loss and flush out. In the case of sufficient nucleation or seeded experiments, condensation on
71 particles is the dominant sink. In absence of nucleation and surface, the other sinks such as wall loss

72 dominate the losses of SOA materials. Therefore, no SOA would be formed without nucleation. This
73 *a priori* tells nothing about the importance of wall losses when SOA is formed.

74 The loss of organic vapors to chamber walls can be important for SOA yield, although in our
75 manuscript we had not corrected for it because it is a challenge to quantify it. In the revised
76 manuscript, we have added a section to estimate the influence of vapor wall loss on SOA yield (Sect.
77 “2.3 Wall loss of organic vapors”) and provided more discussion on vapor wall loss. We found that
78 the influence of vapor wall loss on SOA yield is likely to be significant when surface area
79 concentrations of SOA formed were low in the high NO_x and low SO₂ conditions. Yet, the influence
80 of vapor wall loss is likely to be not significant at the higher surface area concentrations of SOA
81 formed in the low NO_x conditions or high NO_x and high SO₂ conditions.

82 We respectfully disagree with the reviewer’s statement that “previous studies on α -pinene oxidation
83 suggested that SOA yield is independent of particle surface area” without considering each specific
84 study. Although some studies showed that SOA yield from α -pinene oxidation is independent of seed
85 particle surface area (McVay et al., 2016; Nah et al., 2016), a number of studies showed that SOA
86 yield from α -pinene ozonolysis or photooxidation depends on particle surface area. For example, our
87 previous studies clearly showed that SOA yields from α -pinene photooxidation depend on surface
88 area (Sarrafzadeh et al., 2016; Ehn et al., 2014), and Eddingsaas et al. (2012) also showed that in
89 “high NO” conditions, SOA yield from α -pinene photooxidation is much higher with neutral seed
90 than that without seed. The discrepancy in the dependence of SOA yield on particle surface areas in
91 the literature can be attributed to reaction conditions, surface area range and chamber setup. For
92 example, if the reaction produces enough new particles by itself and results in fast particle growth and
93 larger aerosol surface area as the dominant condensational sink for vapors compared to the loss on
94 chamber walls, SOA yield would be less affected by the seed surface area. On the contrary, if VOC
95 oxidation does not induce nucleation by itself, all vapors would be lost onto the chamber walls and
96 SOA yield would be essentially zero.

97 In this study, at high NO_x and low SO₂, the particle number and surface area concentrations were low
98 (peak surface area concentration of $6.8 \times 10^{-7} \text{ cm}^2 \text{ cm}^{-3}$ and particle-to-chamber surface ratio of
99 7.7×10^{-5} for α -pinene), much lower than the aerosol surface area range in the studies by Nah et al.
100 (2016) and McVay et al. (2016) ($\sim 10^{-5} \text{ cm}^2 \text{ cm}^{-3}$ and particle-to-chamber surface ratio of $>4 \times 10^{-5}$). In
101 addition, not only seed particle surface area but also total particle surface area formed during reaction
102 provide condensational sinks to compete with vapor wall loss. At such low particle surface area
103 concentrations, the condensation of vapors on particles had a much longer time scale than that of the
104 wall loss and a large fraction of vapors condensed on chamber walls. Therefore, SOA yield was
105 significantly suppressed due to lack of particle surface area.

106 In the revised manuscript, we have discussed the findings in the literature.

107 “Artificially added seed aerosol has been shown to enhance SOA formation from α -pinene and

108 β -pinene oxidation (Ehn et al., 2014; Sarrafzadeh et al., 2016; Eddingsaas et al., 2012a). In some other
109 studies, it was found that the SOA yield from α -pinene oxidation is independent of initial seed surface
110 area (McVay et al., 2016; Nah et al., 2016). The difference in the literature may be due to the range of
111 total surface area of particles, reaction conditions and chamber setup. For example, the peak
112 particle-to-chamber surface ratio for α -pinene photooxidation in this study was 7.7×10^{-5} at high NO_x
113 and low SO_2 , much lower than the aerosol surface area range in the studies by Nah et al. (2016) and
114 McVay et al. (2016). A lower particle-to-chamber surface ratio can lead to a larger fraction of
115 organics lost on chamber walls. Hence, providing additional particle surface by adding seed particles
116 can increase the condensation of organics on particles and thus increase SOA yield. However, once
117 the surface area is high enough to inhibit condensation of vapors on chamber walls, further
118 enhancement of particle surface will not significantly enhance the yield (Sarrafzadeh et al., 2016).”
119 *The authors explained the effect of SO_2 as 1) inducing new particle formation and providing surface
120 area for vapor condensation, 2) acid-catalyzed particle-phase reactions. I have the same question
121 regarding the first explanation, i.e., what is the role of vapor wall loss (if any), and how does one
122 reconcile this explanation with findings from previous studies? Also, what is the effect of SO_2 on
123 gas-phase chemistry and SOA yield? This is not considered.*

124 **Response:**

125 The role of vapor wall loss can be referred to our response to the comment above. Wall loss of vapors
126 leads to an underestimate of SOA yield. Condensation of vapors onto aerosol particle surface
127 competes with the loss of vapors on chamber walls. Therefore the surface area provided by nucleation
128 and growth of particles in the presence of SO_2 enhanced the SOA yield in this study.

129 Only few studies have investigated the effect of SO_2 on the SOA yield from α -pinene oxidation. More
130 studies investigated the effect of seed aerosol and particle acidity. Kleindienst et al. (2006) attributed
131 the increase of SOA yield in the presence of SO_2 to the formation of H_2SO_4 acidic aerosol. While
132 particle acidity may contribute to the increased SOA yield, especially at high NO_x , the effect of
133 facilitating nucleation and further providing surface area seems to be more important in our study.
134 The importance of the SO_2 via nucleating and providing particle surface depends on the particle
135 surface area in the absence of SO_2 because the competition for the condensation of vapors between
136 particles and wall depends on particle surface area. When VOC oxidation does not form enough new
137 particles and particle surface by itself, the role of SO_2 via nucleating and providing particle surface in
138 enhancing SOA yield is more important. In the revised manuscript, we have added more discussion on
139 this aspect.

140 “Particle acidity may also play a role in affecting the SOA yield in the experiments with high SO_2 .
141 Particle acidity was found to enhance the SOA yield from α -pinene photooxidation at high NO_x
142 (Offenberg et al., 2009) and “high NO ” conditions (Eddingsaas et al., 2012a). Yet, in low NO_x
143 condition, particle acidity was reported to have no significant effect on the SOA yield from α -pinene

144 photooxidation (Eddingsaas et al., 2012a; Han et al., 2016). According to these findings, at low NO_x
145 the enhancement of SOA yield in this study is attributed to the effect of facilitating nucleation and
146 providing more particle surface by SO₂ photooxidation. At high NO_x, the effect in enhancing new
147 particle formation by SO₂ photooxidation seems to be more important, although the effect of particle
148 acidity resulted from SO₂ photooxidation may also play a role.”

149 The effect of SO₂ on gas phase chemistry is not significant in this study because the reaction rate of
150 SO₂ with OH ($\sim 2 \times 10^{-12}$ molecules⁻¹ cm³ s⁻¹) and with RO₂ ($< 10^{-14}$ molecule⁻¹ cm³ s⁻¹) are very low
151 (Lightfoot et al., 1992; Berndt et al., 2015). From the AMS data of SOA formed at high SO₂, we
152 found no significant organic fragments containing sulfur. Also the fragment CH₃SO₂⁺ suggested by
153 Farmer et al. (2010) was not detected in our data. Therefore, we conclude that in our study, the effect
154 of SO₂ on gas phase chemistry of organics and thus further on SOA yield via affecting gas phase
155 chemistry is not important.

156 In the revised manuscript, we have added the following discussion about this point.

157 “SO₂ has been proposed to also affect gas phase chemistry of organics by changing the HO₂/OH or
158 forming SO₃ (Friedman et al., 2016). In this study, the effect of SO₂ on gas phase chemistry of
159 organics was not significant because of the much lower reactivity of SO₂ with OH compared with
160 α-pinene and limonene (Atkinson et al., 2004, 2006; Atkinson and Arey, 2003) and the low OH
161 concentrations (2-3 orders of magnitude lower than those in the study by Friedman et al. (2016)).
162 Moreover, reactions of RO₂ with SO₂ was also not important because the reaction rate constant is very
163 low ($< 10^{-14}$ molecule⁻¹ cm³ s⁻¹) (Lightfoot et al., 1992; Berndt et al., 2015). In addition, from the AMS
164 data of SOA formed at high SO₂ no significant organic fragments containing sulfur were found. Also
165 the fragment CH₃SO₂⁺ from organic sulfate suggested by Farmer et al. (2010) was not detected in our
166 data. The absence of organic sulfate tracers is likely due to the lack of aqueous phase in aerosol
167 particles in this study. Therefore, the influence of SO₂ on gas phase chemistry of organics and further
168 on SOA yield via affecting gas phase chemistry is not important in this study.”

169 *It appears that the SOA yields in high SO₂ experiments might be overestimated by double counting*
170 *the density of ammonium sulfate/ammonium bisulfate in the SOA mass calculation. This is not entirely*
171 *clear.*

172 **Response:**

173 The density of SOA in high SO₂ experiments was not double counted. In the revised manuscript, we
174 have clarified this point. Please also refer to our response to the similar comment below (“detailed
175 comment” 3).

176 “...and their respective density (1.32 g cm⁻³ for organic aerosol from one of our previous studies
177 (Flores et al., 2014) and the literature (Ng et al., 2007) and ~1.77 g cm⁻³ for ammonium
178 sulfate/ammonium bisulfate)...”

179 We would like to note that the SOA yield in this study was derived by adjusting the density of SOA to

180 1 g cm⁻³.

181 *Finally, the authors need to conduct a more careful and accurate comparison with previous studies. It*
182 *was noted that in high SO₂ conditions, their findings that SOA yields are comparable under high NO_x*
183 *and low NO_x conditions are in line with Sarrafzadeh et al. and Eddingsaas et al. I do not think that*
184 *the data in Eddingsaas et al. showed this. SOA yields are also a function of deltaMo (as well as*
185 *various experimental conditions and parameters) and this could play a role, see detailed comment*
186 *below. Also, the α -pinene yields in this study under comparable NO_x/SO₂/OH exposure are much*
187 *lower than Eddingsaas et al.. This is not mentioned and discussed in the manuscript.*

188 **Response:**

189 In the revised manuscript, we have added more discussion to compare this study with previous studies
190 and added a table summarizing previous studies.

191 Please refer to our detailed responses to the specific comments below (“detailed comments” 11d, e).

192 *Major revisions are needed before the manuscript can be published. Specific comments are listed*
193 *below.*

194 *Detailed comments*

195 *1. Line 18-20. This statement is not true for limonene data presented in this study.*

196 **Response:**

197 In the revised manuscript, we have modified this statement. Please refer to our response to a similar
198 comment above (Pg. 1, lines 35-43).

199 *2. Line 79 -81. This sentence seems to imply that previous studies that used higher NO_x and SO₂*
200 *concentrations are not atmospherically relevant. I think these sentences should be revised and*
201 *clarified to more accurately reflect the experimental design and results from previous studies. For*
202 *instance, the use of high levels of NO_x (e.g., from HONO or CH₃ONO) in some studies is to push the*
203 *RO₂ radical fate to the extreme (i.e., RO₂+NO or RO₂+NO₂) to investigate SOA yields and*
204 *composition under such conditions. Thus, the use of high levels of NO_x do not necessarily mean that*
205 *the results are not applicable to ambient conditions.*

206 **Response:**

207 We understand that in some studies high concentration of NO_x is intentionally used to push the RO₂
208 radical fate to an extreme in order to study the SOA yields and composition. Nevertheless,
209 experiments under ambient levels of high NO_x concentrations are more transferable to the ambient
210 anthropogenic-biogenic interaction than the experiments conducted at extremely high NO_x
211 concentrations considering the shortened lifetime of RO₂ and the potential secondary processes as
212 well as the effect of NO_x on OH concentration at extremely high NO_x. In the revised manuscript, we
213 have modified this sentence.

214 “For example, many studies used very high NO_x and SO₂ concentrations (up to several hundreds of
215 ppb). High NO_x can make the RO₂ radical fate dominated by one single pathway (i.e., RO₂+NO or

216 RO₂+NO₂) to investigate SOA yields and composition under such conditions. Yet, the effects of NO_x
217 and SO₂ at concentration ranges for ambient anthropogenic-biogenic interactions (sub ppb to several
218 tens of ppb for NO₂ and SO₂) have seldom been directly addressed.”

219 3. Line 132. Is the organic aerosol density 1.32 g/cm³ from Eddingsaas et al. (2012a)? If this is the
220 case, note that this density used in Eddingsaas et al. is directly taken from the results in Ng et al.
221 (2007), and that this density was obtained in the presence of seeds already. Therefore, it appears that
222 there might be a double counting of the density of ammonium sulfate/ammonium bisulfate in the data
223 presented here?

224 **Response:**

225 Organic aerosol density was based on our previous study (Flores et al., 2014) as well as the study of
226 Eddingsaas et al. (2012) and thus Ng et al. (2007).

227 Although the density of organic aerosol in the study of Ng et al. (2007) was obtained in presence of
228 seed, the contribution of seed aerosol (ammonium sulfate) to particle volume and thus density has
229 been taken into account (Bahreini et al., 2005) . The value reported by Ng et al. (2007) is the density
230 of organic aerosol instead of the density of mixed aerosol. There is no double counting of the density
231 of ammonium sulfate/ammonium bisulfate in our study.

232 4. Line 162. How were OH and O₃ formed in the experiments (under each combination of NO_x/SO₂
233 condition). Please provide more info. Also, please provide typical time profiles of VOC (either
234 *a-pinene* or *limonene*), O₃, OH, NO, NO₂, SO₂ for each combination of NO_x/SO₂ condition. These
235 are important to help the readers obtain a better idea of the reaction pathways/regimes under each
236 condition.

237 **Response:**

238 OH was formed via HONO photolysis, which was produced from a photolytic process on the Teflon
239 chamber wall as we described in the manuscript. The details can be found in a previous study on our
240 chamber (Rohrer et al., 2005). In the revised manuscript, we have added one more sentence to better
241 clarify this point.

242 “OH was formed via HONO photolysis, which was produced from a photolytic process on the Teflon
243 chamber wall (Rohrer et al., 2005).”

244 In addition, in all VOC photooxidation of our study, OH was partly contributed by the recycling
245 reaction of HO₂ with NO. The reaction of HO₂ and RO₂ with NO also produces NO₂, whose
246 photolysis further forms O₃. While the detailed mechanism of O₃ formation is beyond the scope of
247 this study, we have provided a brief description in the revised manuscript.

248 “O₃ was formed in photochemical reactions since NO_x, even in trace amount (<~1 ppbV), was present
249 in this study.”

250 “In the photooxidation of VOC, OH and O₃ often co-exist and both contribute to VOC oxidation
251 because O₃ formation in chamber studies is often unavoidable during photochemical reactions of

252 VOC even in the presence of trace amount of NO_x.”

253 In the revised manuscript, we have provided time profiles of VOC, O₃, OH, NO, and NO₂ in low NO_x
254 and high NO_x conditions for α-pinene and limonene (Fig. S5). Time profiles of these species at high
255 SO₂ were similar to those at low SO₂ because SO₂ had little effect on gas phase chemistry due to its
256 low reactivity of SO₂ with OH and RO₂, and thus are not further shown. SO₂ time series are shown
257 separately (Fig. S2).

258 5. Line 163. Was all the VOC reacted in each experiment?

259 **Response:**

260 All the VOC precursor was consumed in the experiments of this study except for one experiment
261 where small amount of VOC (~10%) was still left at the end of the reaction due to the cloudy
262 condition and thus lower photolysis rates.

263 6. Lin 163. There is no “typical” experiment in this study, as each experiment was conducted under a
264 different NO_x/SO₂ condition. Please clearly state that this is only for low NO_x condition. Also, what
265 about high NO_x condition? Was it exclusively OH reaction? Please also specify clearly.

266 **Response:**

267 The relative dominance of OH oxidation over ozonolysis (as shown in Fig. 6) is similar in both the
268 low NO_x and high NO_x conditions. At high NO_x, OH was often higher and meanwhile more O₃ was
269 also produced.

270 In the revised manuscript, we have modified this sentence as follows.

271 “For all the experiment in this study, the VOC loss was dominated by OH oxidation over ozonolysis
272 (see Fig. S6 as an example). The relative importance of the reaction of OH and O₃ with monoterpenes
273 was similar in the low NO_x and high NO_x experiments. At high NO_x, OH was often higher while more
274 O₃ was also produced.”

275 7. Line 173 – 177. Here, under low NO_x condition, RO₂+NO dominates throughout the entire
276 experiment (RO₂+HO₂ only contributes to 40% at most).

277 a. These sentences clearly demonstrate the shortcomings of classifying the experiments as low NO_x vs.
278 high NO_x as discussed in Wennberg et al. (IGAC news, 2103). I suggest the authors to characterize
279 reactions conditions by explicitly stating the RO₂ fates, rather than as low vs. high NO_x.

280 **Response:**

281 We had clearly defined our low NO_x and high NO_x conditions using the RO₂ fate in our study.
282 Therefore, we respectfully do not think that using the terms “low NO_x” and “high NO_x” caused
283 ambiguity as long as we define them clearly.

284 In the revised manuscript, we have added the following sentence to emphasize the RO₂ fates under
285 different NO_x conditions.

286 “Note that the RO₂ fate in the low and high NO_x conditions quantified here are further used in the
287 discussion below since the information of RO₂ fate is important for data interpretation of experiments

288 conducted at different NO_x levels (Wennberg, 2013).”

289 *b. It is stated that under low NO_x conditions, in the beginning of the experiment, a trace amount of*
290 *NO is formed from photolysis of HONO from the chamber wall. Is this just in the beginning of the*
291 *experiment, or there is a continuous NO source from HONO photolysis throughout the entire*
292 *experiment? Please specify.*

293 **Response:**

294 There is a continuous NO source from HONO photolysis. In the revised manuscript, we have clarified
295 this as follows.

296 “The trace amount of NO (up to ~0.2 ppbV) was from the photolysis of HONO, which was
297 continuously produced from a photolytic process on chamber walls throughout an experiment (Rohrer
298 et al., 2005).”

299 *8. Line 189. The authors attributed the lower particle number concentration and nucleation rate at*
300 *high NO_x to the decreasing RO₂+RO₂ reaction in the presence of NO_x. However, in line 182, the*
301 *authors noted that RO₂+RO₂ reaction is negligible in this study to start with. Please reconcile these*
302 *seemingly contradictory statements. Also, can be suppressed nucleation under high NO_x due to the*
303 *higher volatility of organic nitrates as compared to peroxides (from RO₂+HO₂)?*

304 **Response:**

305 The compounds responsible for nucleation only account for a very small fraction of RO₂ reaction
306 products. Although the contribution of RO₂+RO₂ reaction to the total RO₂ loss is negligible, it can
307 contribute a lot to the compounds responsible for nucleation because RO₂+RO₂ reactions form dimers,
308 which have high molecular weight and extremely low volatility (Ehn et al., 2014; Kirkby et al., 2016).
309 Generally, organic nitrates are not expected to be the main compounds responsible for nucleation
310 since their volatility is not low enough to nucleate, nor primary organic peroxides (from
311 RO₂(C₁₀)+HO₂). Therefore, although under high NO_x more organic nitrates were found, organic
312 nitrates are unlikely to be the reason for the suppressed nucleation.

313 *9. Line 205-206. There is nucleation (from organics) in the presence of NO_x as shown in Fig. 4. In*
314 *this sense, “absence of nucleation” here is a bit confusing. Perhaps would be clearer to say “absence*
315 *of seed particles”.*

316 **Response:**

317 In the revised manuscript, we have modified “absence of nucleation” as follows.

318 “Because NO_x suppressed new particle formation, the suppression of the SOA yield could be
319 attributed to the lack of new particles as seed and thus the lack of condensational sink, or to the
320 decrease of condensable organic materials.”

321 *10. Line 211 and Figure 3. The author concluded that the suppression effect of NO_x on SOA yields*
322 *was mainly due to suppression of nucleation, i.e., to the absence of particle surface as condensation*
323 *sink. Many critical aspects are not discussed, making this conclusion not well-justified and*

324 *well-supported.*

325 *a. If the absence of seed particle surface area is the reason for the low yield under high NO_x*
326 *condition (at low SO₂), this will point to the importance of loss of semivolatile species via chemical*
327 *reactions or chamber wall loss (Kroll et al., ES&T, 2007). However, the effect of vapor wall loss in*
328 *not considered in this study. Zhang et al. (PNAS, 2014) first systematically investigated the effects of*
329 *particle surface area and vapor wall loss on SOA yields. For α -pinene photooxidation and ozonolysis*
330 *specifically, it was found that SOA yields are largely independent of seed surface area (McVay et al.,*
331 *2016, Nah et al., ACP, 2016; Nah et al., ACP, 2017). Therefore, taken all these together, it is not*
332 *clear how the absence of particle surface area can explain the suppressed SOA yields under high NO_x*
333 *condition in this study.*

334 **Response:**

335 Please refer to our response to a similar comment above (Pg. 2, lines 58-66).

336 *b. The authors dismissed the “decrease of condensable organic materials” in high NO_x conditions as*
337 *an explanation for the observed decrease in yield. Why? If more volatile organic nitrates are formed*
338 *in high NO_x conditions, why can’t this be an explanation for the suppressed SOA yield? For limonene*
339 *data (Line 213-218), the authors appeared to embrace the role of volatility of oxidation products.*

340 **Response:**

341 In our study, when new particle formation was already enhanced by added SO₂, the SOA yield at high
342 NO_x was comparable to that at low NO_x for α -pinene and the difference in SOA yield between high
343 NO_x and low NO_x was much smaller (Fig. 3a). If the organic materials such as organic nitrate formed
344 in high NO_x conditions were more volatile, the SOA yield in high NO_x should be low regardless of
345 SO₂ concentration unless in addition to that, SO₂ enhanced the SOA yield at high NO_x via the
346 influence other than surface area effect, e.g., acidity effect. Organic nitrates formed at high NO_x was
347 proposed to be more volatile (Presto et al., 2005; Kroll et al., 2006). However, many organic nitrate
348 formed in our study is highly oxidized organic molecule (HOMs) containing multi-functional groups
349 besides nitrate (C₇₋₁₀H₉₋₁₅NO₈₋₁₅). The compounds are expected to have low volatility and they were
350 found to have an uptake coefficient on particles of ~1 (Pullinen et al. in preparation). A recent study
351 also implied that organic nitrate may have low volatility (Hakkinen et al., 2012). Therefore, the
352 suppressing effect of NO_x on SOA yield was more likely due to suppressed nucleation, i.e., to lack of
353 particle surface as condensational sinks.

354 In the revised manuscript, we have elaborated our discussion.

355 “This finding can be attributed to two possible explanations. Firstly, NO_x did not significantly
356 suppress the formation of low volatile condensable organic materials, although NO_x obviously
357 suppressed the formation of products for nucleation. Secondly, NO_x did suppress the formation of
358 low-volatility condensable organic materials via forming potentially more volatile compounds and in
359 addition to that, the suppressed formation of condensable organic materials was compensated by the

360 presence of SO₂, resulting in comparable SOA yield. Organic nitrates are a group of compounds
361 formed at high NO_x, which have been proposed to be more volatile (Presto et al., 2005; Kroll et al.,
362 2006). However, many organic nitrates formed by photooxidation in this study were highly oxidized
363 organic molecules (HOMs) containing multi-functional groups besides nitrate group
364 (C₇₋₁₀H₉₋₁₅NO₈₋₁₅). These compounds are expected to have low volatility and they are found to have an
365 uptake coefficient on particles of ~1 (Pullinen et al., in preparation). Therefore, the suppressing
366 effect of NO_x on SOA yield was **mostly likely** due to suppressed nucleation, i.e., **the lack** of particle
367 surface as condensational sink.”

368 For limonene data, please refer to our response to the comment below (“detailed comments” 10c).

369 *c. Line 217. How does the different range in VOC/NO_x for α -pinene and limonene experiments*
370 *explain the differences in yields in high SO₂ conditions? Please elaborate and explain clearly.*

371 **Response:**

372 The cause of the difference between the α -pinene and limonene cases is unknown for the moment and
373 it would be pure speculation when discussing reasons for this difference. Therefore, in the revised
374 manuscript, we only state that the reason is unknown so far and as possible explanation, we note that
375 the average volatility of limonene oxidation products may be higher at higher NO_x.

376 In the revised manuscript, we have revised this sentence as follows.

377 “**The cause of this difference is currently unknown. Our data of SOA yield suggest that the products**
378 **formed from limonene oxidation at high NO_x seemed to have higher average volatility than that at low**
379 **NO_x.**”

380 *11. Line 225-237. Comparisons with previous studies. Many critical details are not considered and*
381 *discussed. I think the authors jumped to the conclusion on whether their study agree/disagree with*
382 *previous studies too quickly.*

383 *a. Line 225. This sentence is only true for α -pinene data in this study, but not for limonene. Please*
384 *state clearly.*

385 **Response:**

386 In the revised manuscript, we have revised this sentence as follows.

387 “Our finding that **the difference in SOA yield between high NO_x and low NO_x conditions was highly**
388 **reduced at high SO₂** is also in line with the findings of some previous studies using seed aerosols
389 (Sarrafzadeh et al., 2016; Eddingsaas et al., 2012a).”

390 *b. Seed particles were generated via SO₂ oxidation in this study (for high SO₂ experiments).*

391 *Previous studies directly injected seeds into the chamber. In comparing SOA yields, the author should*
392 *also consider the role of gas-phase chemistry and particle phase chemistry. For instance, what about*
393 *the reaction of SO₂ and criegee intermediates? What about the effect of particle acidity on*
394 *particle-phase reactions (in this study vs. previous studies)? Please discuss.*

395 **Response:**

396 In our study, the reaction of SO₂ with Criegee intermediates was not important to the formation of
397 oxidized organics and SOA formation for the following reasons. 1) The reaction of VOC with O₃ only
398 contributed to a small fraction of VOC loss in this study and thus formation of Criegee intermediates
399 was not significant. 2) At the water vapor concentration of this study, water may compete for Criegee
400 intermediates with SO₂ to a large extent.

401 Particle acidity may affect the SOA yield via acid-catalyzed reactions, as we had discussed in our
402 manuscript. In the revised manuscript, we have elaborated this discussion by comparing with previous
403 studies as follows.

404 “Particle acidity may also play a role in affecting the SOA yield in the experiments with high SO₂.
405 Particle acidity was found to enhance the SOA yield from α-pinene photooxidation at high NO_x
406 (Offenberg et al., 2009) and “high NO” conditions (Eddingsaas et al., 2012a). Yet, in low NO_x
407 condition, particle acidity was reported to have no significant effect on the SOA yield from α-pinene
408 photooxidation (Eddingsaas et al., 2012a; Han et al., 2016). According to these findings, at low NO_x
409 the enhancement of SOA yield in this study is attributed to the effect of facilitating nucleation and
410 providing more particle surface by SO₂ photooxidation. At high NO_x, the effect in enhancing new
411 particle formation by SO₂ photooxidation seems to be more important, although the effect of particle
412 acidity resulted from SO₂ photooxidation may also play a role.”

413 *c. The experiments in this study were conducted in the presence of humidity but previous studies were*
414 *mostly conducted under dry conditions. RH can affect gas-phase and particle-phase chemistry, and*
415 *subsequently SOA yields.*

416 **Response:**

417 We agree with the reviewer that RH can affect gas-phase and particle-phase chemistry and thus may
418 also affect SOA yield. Because humidity is ubiquitous in the real atmosphere, we conducted our
419 experiments in the presence of humidity in order to better represent ambient conditions. In the revised
420 manuscript, we have emphasized this point.

421 “RH of this study is different from many previous studies, which often used very low RH (<10%).”

422 *d. The authors noted that the finding that SOA yields at high NO_x is comparable to that at low NO_x in*
423 *high SO₂ conditions is in line with findings in Sarrafzadeh et al. (2016) and Eddingsaas et al. (2012a).*
424 *I do not think that the data in Eddingsaas et al. showed that “in presence of seed aerosol, the*
425 *difference in the SOA yield between low and high NO_x is much reduced”. SOA yield is also a function*
426 *of deltaMo. Considering the data in Table 1 of Eddingsaas et al., the difference in yields between low*
427 *and high NO experiments for nucleation is 19%, and for seeded experiments are 15% and 10%.*

428 *However, the difference in deltaMo for the nucleation experiments is also the largest and this will*
429 *play a role in the yield difference.*

430 **Response:**

431 As the reviewer noticed, the data in Eddingsaas et al. (2012) showed that in the absence of seed

432 aerosol SOA yield at low NO_x is 2.5 times higher than that at high NO_x, while in the presence of seed
433 aerosol SOA yield at low NO_x is only 1 and 0.6 times higher than that at high NO_x for neutral seed
434 and acid seed, respectively. Therefore, the data in Eddingsaas et al. (2012) did show that “in presence
435 of seed aerosol, the difference in the SOA yield between low and high NO_x is much reduced”.

436 We noticed that in the data given by Eddingsaas et al. (2012) the difference in deltaM₀ for nucleation
437 experiments is the largest, which plays a role in SOA yield. However, the large difference in deltaM₀
438 between high NO_x and low NO_x cases is because the deltaM₀ at high NO_x is the lowest in absence of
439 seed, much lower than deltaM₀ in presence of seed when other conditions are largely the same. The
440 higher deltaM₀ and the smaller difference in deltaM₀ between low NO_x and high NO_x in presence of
441 seed also originated from the seed aerosol since other conditions were kept constant. This result
442 agrees with our finding that “in presence of seed aerosol, the difference in the SOA yield between low
443 and high NO_x is much reduced”.

444 *e. The SOA yields in this study are much lower than previous studies, why? Considering the low NO_x*
445 *low SO₂ experiment, with OH dose of 1e11 molecules cm⁻³ s, the yield in this study is 7%. However,*
446 *the corresponding yield in Eddingsaas et al. is > 30% (Figure 3 of Eddingsaas et al.).*

447 **Response:**

448 The difference in SOA yield between this study and the study by Eddingsaas et al. (2012) can be
449 explained by several reasons. Firstly, SOA yield in this study was calculated using a density of 1 g
450 cm⁻³ while in Eddingsaas et al. (2012) SOA yield was calculated using a density of 1.32 g cm⁻³.
451 Secondly, reaction conditions such as VOC concentrations, NO_x concentrations, and OH source and
452 concentrations of our study at low NO_x are different from those in Eddingsaas et al. (2012). For
453 example, in our study at low NO_x, RO₂+NO account for a large fraction of RO₂ loss while in
454 Eddingsaas et al. (2012) RO₂+HO₂ is the dominant pathway of RO₂ loss. These differences in reaction
455 conditions may affect SOA yield. Thirdly, the organic aerosol concentration of this study is much
456 lower than that in Eddingsaas et al. (2012). Fourthly, the total particle surface area may be different
457 from Eddingsaas et al. (2012) (the data are not available to compare with), which can also affect the
458 measured SOA yield. Also note that the exceptionally high SOA yield at lower α-pinene
459 concentration is an “outlier” to other data in Eddingsaas et al. (2012) and could not be explained by
460 the authors.

461 In the revised manuscript, we have added the following discussion regarding the comparison of SOA
462 yield from α-pinene photooxidation with the literature.

463 “The reaction conditions of this study often differ from those described in the literature (see Table
464 S2).

465 The difference in these conditions can result in both different apparent dependence on specific
466 parameters and the varied SOA yield. For example, SOA yield from α-pinene photooxidation at low
467 NO_x in this study appeared to be much lower than that in Eddingsaas et al. (2012a). The difference

468 between the SOA yield in this study and some of previous studies and between the values in the
469 literature can be attributed to several reasons: 1) RO₂ fates may be different. For example, in our study
470 at low NO_x, RO₂+NO account for a large fraction of RO₂ loss while in Eddingsaas et al. (2012a)
471 RO₂+HO₂ is the dominant pathway of RO₂ loss. This difference in RO₂ fates may affect oxidation
472 products distribution. 2) The organic aerosol loading of this study is much lower than that some of
473 previous studies (e.g., Eddingsaas et al. (2012a)) (see Fig. S9). SOA yields in this study were also
474 plotted versus organic aerosol loading to better compare with previous studies (Fig. S8 and S9). 3)
475 The total particle surface area in this study may also differ from previous studies, which may
476 influence the apparent SOA yield due to vapor wall loss (the total particle surface area is often not
477 reported in many previous studies to compare with). 4) RH of this study is different from many
478 previous studies, which often used very low RH (<10%). It is important to emphasize that reaction
479 conditions including the NO_x as well as SO₂ concentration range and RH in this study were chosen to
480 be relevant to the anthropogenic-biogenic interactions in the ambient atmosphere. In addition,
481 difference in the organic aerosol density used in yield calculation should be taken into account. In this
482 study, SOA yield was derived using a density of 1 g cm⁻³ to better compare with many previous
483 studies (e.g., (Henry et al., 2012)), while in some other studies SOA yield was derived using different
484 density (e.g., 1.32 g cm⁻³ in Eddingsaas et al. (2012a)).”

485 In the revised manuscript, we have also added a figure plotting SOA yield as a function of organic
486 aerosol mass loading (Fig. S8).

487 *12. Line 238 onwards, effect of SO2.*

488 *a. One of the proposed reasons to explain the effect of SO2 is that it induces nucleation and provides*
489 *more particle surface area for condensation. Again, if this is the case, it will point to the importance*
490 *of loss of organic vapors to chamber walls, though previous studies suggested that this process does*
491 *not effect SOA yields from α-pinene oxidations to a large extent. With this, it is not clear if this is*
492 *indeed a reason for the observed SO2 effect. Please explain.*

493 **Response:**

494 Please refer to our response to the similar comments above (Pg. 2, lines 58-66; Pg. 4, lines 119-123)
495 and corresponding revisions to the manuscript.

496 *b. Line 258. Is “counterbalance” the appropriate word? If the suppression effect of NOx is*
497 *counterbalanced by the enhancement effect SO2, in going from “low NOx low SO2” to “high NOx*
498 *high SO2” one shall not observe change in SOA yields? Also, note that the limonene data showed*
499 *very different trends comparing to the α-pinene data. This needs to be accurately and clearly stated.*

500 **Response:**

501 We have changed “counterbalance” to “compensated”. The revised sentence is as follows.

502 “The presence of high SO₂ enhanced the SOA mass yield at high NO_x conditions, which was even
503 comparable with the SOA yield at low NO_x for α-pinene oxidation. This finding indicates that the

504 suppressing effect of NO_x on SOA mass formation was compensated to large extent by the presence
505 of SO₂.”

506 As for the difference in the limonene and α -pinene data, please refer to our response to the similar
507 comments above (Pg. 1, lines 35-43) and our corresponding revisions to the manuscript.

508 *13. Line 315-318. This explanation is a stretch and not well-justified. There is extensive fragmentation*
509 *in the AMS and so the H/C ratios of oxidation product molecules do not necessarily translate to the*
510 *H/C ratios measured. As shown in Chhabra et al., not all experiments conducted under low NO_x*
511 *condition have higher H/C ratios.*

512 **Response:**

513 In our opinion, the reaction pathway of RO₂ at least provides a likely explanation for the difference in
514 H/C at different NO_x. Although there is extensive fragmentation in the AMS, H/C measured generally
515 reflects the H/C ratios of the overall compounds. In addition, as we discussed, Chhabra et al. (2011)’s
516 data also show that for α -pinene photooxidation, SOA formed at high NO_x generally has lower H/C,
517 consistent with our study. Admittedly, the RO₂ reaction pathway is not the only factor affecting H/C
518 and O/C. Other factors such as VOC identity, oxidants, and reaction mechanisms, including various
519 functionalization, fragmentation, and oligomerization in both the gas phase and particle phase also
520 play important roles in the chemical composition and thus H/C and O/C. We did not intend to apply
521 our explanation here to all other reaction systems.

522 *14. Line 262 onwards. Did the ratio of nitrate mass concentration to organics mass change over*
523 *time?*

524 **Response:**

525 The ratio of nitrate mass concentration to organics gradually decreased in the beginning of the
526 reactions (2-3 h) and then leveled off. Also, note that in the very beginning of a reaction, the data have
527 large uncertainties due to the low concentration of nitrate and organics. In our manuscript, the average
528 ratios were used to compare different experiments. In the revised manuscript, we have clearly
529 described this in the captions of Fig. 4.

530 “The average ratios of nitrate to organics during the reaction are shown and error bars indicate the
531 standard deviations.”

532 Minor comments

533 *1. Line 72. Why “in contrast”?*

534 **Response:**

535 The finding here is different from those in the studies discussed before. In the revised manuscript, we
536 have modified this sentence as follows.

537 “In contrast, Eddingsaas et al. (2012a) found that particle yield increases with aerosol acidity only in
538 “high NO” condition (NO_x 800 ppb, α -pinene: 20-52 ppb), but is independent of the presence of seed
539 aerosol or aerosol acidity in both “high NO₂” condition (NO_x 800 ppb)” and low NO_x (NO_x lower than

540 the detection limit of the NO_x analyzer).”

541 2. Line 84 citation. There are more studies on OH oxidations of α -pinene and they should also be
542 cited here (for example, some of the studies cited in page 2).

543 **Response:**

544 In the revised manuscript, we have added more studies on OH oxidation of α -pinene. However, we
545 would like to note here that many studies on α -pinene photooxidation did not quantitatively
546 distinguish the contributions of oxidation by OH and by O₃.

547 3. Line 125. “mass” should be “volume”? SMPS measures volume concentration.

548 **Response:**

549 We have changed the “mass” to “volume” in the revised manuscript.

550 4. Line 126. Delete “with”.

551 **Response:**

552 Done.

553 5. Line 256. Sentence not clear.

554 **Response:**

555 This sentence has been revised as follows.

556 “The presence of high SO₂ enhanced the SOA mass yield at high NO_x conditions, which was even
557 comparable with the SOA yield at low NO_x for α -pinene oxidation. This finding indicates that the
558 suppressing effect of NO_x on SOA mass formation was compensated to large extent by the presence
559 of SO₂.”

560 6. Figure 1 caption should specify the SO₂ condition.

561 **Response:**

562 Done.

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Effects of NO_x and SO₂ on the Secondary Organic Aerosol Formation from Photooxidation of α -pinene and Limonene

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Abstract

Anthropogenic emissions such as NO_x and SO₂ influence the biogenic secondary organic aerosol (SOA) formation, but detailed mechanisms and effects are still elusive. We studied the effects of NO_x and SO₂ on the SOA formation from the photooxidation of α -pinene and limonene at ambient relevant NO_x and SO₂ concentrations (NO_x: < 1 ppb to 20 ppb, SO₂: <0.05 ppb to 15 ppb). In these experiments, monoterpene oxidation was dominated by OH oxidation. We found that SO₂ induced nucleation and enhanced SOA mass formation. NO_x strongly suppressed not only new particle formation but also SOA mass yield. However, in the presence of SO₂ which induced high number concentration of particles after oxidation to H₂SO₄, the suppression of the mass yield of SOA at highby NO_x was comparable to that at low NO_x completely or partly compensated. This indicates that the suppression of SOA yield by NO_x was mainlylargely due to the suppressed new particle formation, leading to a lack of particle surface for the organics to condense on. By compensating for the suppressing effect on nucleation of NO_x, SO₂ also compensated for the suppressing effect on SOA yield. Aerosol mass spectrometer data show that increasing NO_x enhanced nitrate formation. The majority of the nitrate was organic nitrate (57%-77%), even in low NO_x conditions (<~1 ppb). Organic nitrate contributed 7%-26% of total organics assuming a molecular weight of 200 g/mol. SOA from α -pinene photooxidation at high NO_x had generally lower hydrogen to carbon ratio (H/C), compared with-atto low NO_x. The NO_x dependence of the chemical composition can be attributed to the NO_x dependence of the branching ratio of the RO₂ loss reactions, leading to lower fraction of organic hydroperoxides and higher fractions of organic nitrates at high NO_x. While NO_x suppressed new particle formation and SOA mass formation, SO₂ can compensate for such effects, and the combining effect of SO₂ and NO_x may have important influence on SOA formation affected by interactions of biogenic volatile organic compounds (VOC) with anthropogenic emissions.

34 1 Introduction

35 Secondary organic aerosol (SOA), ~~as an important class of atmospheric aerosol, have important~~ have
36 significant impacts on air quality, human health and climate change (Hallquist et al., 2009; Kanakidou et al.,
37 2005; Jimenez et al., 2009; Zhang et al., 2011). SOA mainly originates from biogenic volatile organic compounds
38 (VOC) emitted by terrestrial vegetation (Hallquist et al., 2009). Once emitted into the atmosphere, biogenic VOC
39 can undergo reactions with atmospheric oxidants including OH, O₃ and NO₃, and form SOA. When an air mass
40 enriched in biogenic VOC is transported over an area with substantial anthropogenic emissions or vice versa, the
41 reaction behavior of VOC and SOA formation can be altered due to the interactions of biogenic VOC with
42 anthropogenic emissions such as NO_x, SO₂, anthropogenic aerosol and anthropogenic VOC. A number of field
43 studies have highlighted the ~~important~~ role of the anthropogenic-biogenic interactions in SOA formation (de
44 Gouw et al., 2005; Goldstein et al., 2009; Hoyle et al., 2011; Worton et al., 2011; Glasius et al., 2011; Xu et al.,
45 2015a; Shilling et al., 2012), which can induce an “anthropogenic enhancement” effect on SOA formation.

46 Among biogenic VOC, monoterpenes are important contributors to biogenic SOA due to their high emission
47 rates, high reactivity and relative high SOA yield compared to isoprene (Guenther et al., 1995; Guenther et al.,
48 2012; Chung and Seinfeld, 2002; Pandis et al., 1991; Griffin et al., 1999; Hoffmann et al., 1997; Zhao et al.,
49 2015; Carlton et al., 2009). The anthropogenic modulation of the SOA formation from monoterpene can have
50 important impacts on regional and global biogenic SOA budget (Spracklen et al., 2011). The influence of various
51 anthropogenic pollutants on SOA formation of monoterpene have been investigated by a number of laboratory
52 studies (Sarrafzadeh et al., 2016; Zhao et al., 2016; Flores et al., 2014; Emanuelsson et al., 2013; Eddingsaas et
53 al., 2012a; Offenberg et al., 2009; Kleindienst et al., 2006; Presto et al., 2005; Ng et al., 2007; Zhang et al., 1992;
54 Pandis et al., 1991; Draper et al., 2015; Han et al., 2016). In particular, NO_x and SO₂ have been shown to affect
55 SOA formation- from monoterpene.

56 NO_x changes the fate of RO₂ radical formed in VOC oxidation and therefore can change reaction product
57 distribution and aerosol formation. At low NO_x, RO₂ mainly react with HO₂, forming organic hydroperoxides. At
58 high NO_x, RO₂ mainly react with NO₂ forming organic nitrate (Hallquist et al., 2009; Ziemann and Atkinson,
59 2012; Finlayson-Pitts and Pitts Jr., 1999). Some studies found that the SOA yield from α -pinene is higher at
60 lower NO_x concentration for ozonolysis (Presto et al., 2005) and photooxidation (~~Ng et al., 2007; Eddingsaas et~~
61 ~~al., 2012a; Han et al., 2016~~)(Ng et al., 2007; Eddingsaas et al., 2012a; Han et al., 2016; Stirnweis et al., 2017).
62 The decrease of SOA yield with increasing NO_x was proposed to be due to the formation of more volatile
63 products like organic nitrate under high NO_x conditions (Presto et al., 2005). In contrast, a recent study found that
64 the suppressing effect of NO_x is in large part attributed to the effect of NO_x on OH concentration for the SOA
65 from β -pinene oxidation, and after eliminating the effect of NO_x on OH concentration, SOA yield only varies by
66 20-30% (Sarrafzadeh et al., 2016). Beside the effect of NO_x on SOA yield, NO_x has been found to suppress the
67 new particle formation from VOC directly emitted by Mediterranean trees (mainly monoterpenes) (Wildt et al.,
68 2014) and β -pinene (Sarrafzadeh et al., 2016), thereby reducing condensational sink present during high NO_x
69 experiments.

70 Regarding the effect of SO₂, the SOA yield of α -pinene photooxidation was found to increase with SO₂
71 concentration at high NO_x concentrations (SO₂: 0-252 ppb, NO_x: 242-543 ppb, α -pinene: 178-255 ppb)

72 (Kleindienst et al., 2006) and the increase is attributed to the formation of H₂SO₄ acidic aerosol. Acidity of seed
73 aerosol was also found to enhance particle yield of α -pinene at high NO_x (Offenberg et al. (2009): NO_x 100-120
74 ppb, α -pinene 69-160 ppb; Han et al. (2016): initial NO ~70 ppb, α -pinene 14-18 ppb). In contrast, Eddingsaas et
75 al. (2012a) found that particle yield increases with aerosol acidity ~~at only in~~ “high NO” condition (NO_x 800 ppb,
76 α -pinene: 20-52 ppb), but is independent of the presence of seed aerosol or aerosol acidity ~~at in~~ both “high NO₂”
77 condition (NO_x 800 ppb)” and low NO_x (NO_x lower than the detection limit of the NO_x analyzer) and “high NO₂”
78 condition (NO_x 800 ppb). Similarly, at low NO_x (initial NO <0.3 ppb, α -pinene ~20 ppb), Han et al. (2016) found
79 that the acidity of seed has no significant effect on SOA yield from α -pinene photooxidation. In addition, SO₂
80 was found to influence the gas phase oxidation products from α -pinene and β -pinene photooxidation, which is
81 possibly due to the change in OH/HO₂ ratio caused by SO₂ oxidation or SO₃ directly reacting with organic
82 molecules (Friedman et al., 2016).

83 While these studies have provided ~~important~~valuable insights into the effects of NO_x and SO₂ on SOA
84 formation, a number of questions still remain elusive. For example, many studies used very high NO_x and SO₂
85 concentrations (up to several hundreds of ppb~~);~~). High NO_x can make the RO₂ radical fate dominated by one
86 single pathway (i.e., RO₂+NO or RO₂+NO₂) to investigate SOA yields and composition under such conditions.
87 Yet, the effects of NO_x and SO₂ at ~~concentrations relevant to~~concentration ranges for ambient anthropogenic-
88 biogenic interactions (sub ppb to several tens of ppb for NO₂ and SO₂) ~~are unclear~~have seldom been directly
89 addressed. Moreover, many previous studies on the SOA formation from monoterpene oxidation focus on
90 ozonolysis or do not distinguish ~~the~~ OH oxidation and ozonolysis in photooxidation, and only a few studies on
91 OH oxidation have been conducted (~~Eddingsaas et al., 2012a; Zhao et al., 2015~~)(Eddingsaas et al., 2012a; Zhao et
92 al., 2015; McVay et al., 2016; Sarrafzadeh et al., 2016; Henry et al., 2012; Ng et al., 2007). More importantly,
93 studies that investigated the combined effects of NO_x and SO₂ are scarce, although they are often co-emitted from
94 anthropogenic sources. According to previous studies, NO_x ~~mainly has~~can have a suppressing effect on SOA
95 formation while SO₂ ~~mainly has~~can have an enhancing effect. NO_x and SO₂ might have counteracting ~~effect~~ or a
96 synergistic effects in SOA formation in the ambient atmosphere.

97 In this study, we investigated the effects of NO_x, SO₂ and their combining effects on SOA formation from the
98 photooxidation of α -pinene and limonene. α -pinene and limonene are two important monoterpenes with high
99 emission rates among monoterpenes (Guenther et al., 2012). OH oxidation dominated over ozonolysis in the
100 monoterpene oxidation in this study as determined by measured OH and O₃ concentrations. The relative
101 contributions of RO₂ loss reactions at low NO_x and higher NO_x were ~~also~~ quantified using measured HO₂, RO₂,
102 and NO concentrations. The effects on new particle formation, SOA yield and aerosol chemical composition
103 were examined. We used ambient relevant NO_x and SO₂ concentrations so that the results can shed lights on the
104 mechanisms of interactions of biogenic VOC with anthropogenic emissions in the real atmosphere.

105 2 Experimental

106 2.1 Experimental setup and instrumentation

107 The experiments were performed in the SAPHIR chamber (Simulation of Atmospheric PHotochemistry In a
108 large Reaction chamber) at Forschungszentrum Jülich, Germany. The details of the chamber have been described
109 before (Rohrer et al., 2005; Zhao et al., 2015). Briefly, it is a 270 m³ Teflon chamber using natural sunlight for
110 illumination. It is equipped with a louvre system to switch between light and dark conditions. The physical
111 parameters for chamber running such as temperature and relative humidity were recorded. The solar irradiation
112 was characterized and the photolysis frequency was derived (Bohn et al., 2005; Bohn and Zilken, 2005).

113 Gas and particles phase species were characterized using various instruments. OH, HO₂ and RO₂
114 concentrations were measured using a laser induced fluorescence (LIF) system with details described by Fuchs et
115 al. (2012). OH was formed via HONO photolysis, which was produced from a photolytic process on the Teflon
116 chamber wall (Rohrer et al., 2005). From OH concentration, OH dose, the integral of OH concentration over
117 time, was calculated in order to better compare experiments with different OH levels. For example, experiments
118 at high NO_x in this study generally had higher OH concentrations due to the faster OH production by recycling of
119 HO₂• and RO₂• to OH. The VOC were characterized using a Proton Transfer Reaction Time-of-Flight Mass
120 Spectrometer (PTR-ToF-MS) and Gas Chromatography-Mass spectrometer (GC-MS). NO_x, O₃ and SO₂
121 concentrations were characterized using a NO_x analyzer (ECO PHYSICS TR480), an O₃ analyzer (ANSYCO,
122 model O341M), and an SO₂ analyzer (Thermo Systems 43i), respectively. O₃ was formed in photochemical
123 reactions since NO_x, even in trace amount (<~1 ppbV), was present in this study. More details of these
124 instrumentation are described before (Zhao et al., 2015).

125 The number and size distribution of particles were measured using a condensation particle counter (CPC,
126 TSI, model 3786) and a scanning mobility particle sizer (SMPS, TSI, DMA 3081/CPC 3785). From particle
127 number measurement, the nucleation rate ($J_{2.5}$) was derived from the number concentration of particles larger than
128 2.5 nm as measured by CPC. Particle chemical composition was measured using a High-Resolution Time-of-
129 Flight Aerosol Mass Spectrometer (HR-ToF-AMS, Aerodyne Research Inc.). From the AMS data, oxygen to
130 carbon ratio (O/C), hydrogen to carbon ratio (H/C), and nitrogen to carbon ratio (N/C) were derived using a
131 method derived in the literature (Aiken et al., 2007; Aiken et al., 2008). An update procedure to determine the
132 elemental composition is reported by Canagaratna et al. (2015), showing the O/C and H/C derived from the
133 method of Aiken et al. (2008) ~~may be underestimated. Since only relative difference in elemental composition of~~
134 ~~SOA is studied here, different methods should not affect the relative trend in elemental composition. The~~
135 ~~fractional contribution of organics in the signals at $m/z=44$ and $m/z=43$ to total organics (f_{44} and f_{43} , respectively)~~
136 ~~were also derived. may be underestimated. The H/C and O/C were also derived using the newer approach by~~
137 Canagaratna et al. (2015) and compared with the data derived from the Aiken et al. (2007) method. The H/C
138 values derived using the Canagaratna et al. (2015) method strongly correlated with the values derived using
139 Aiken et al. (2007) method (Fig. S1) and just increased by 27% as suggested by Canagaratna et al. (2015).
140 Similar results were found for O/C and there was just a difference of 11% in O/C. Since only relative difference
141 in elemental composition of SOA is studied here, only the data derived using Aiken et al. (2007) method are

142 ~~shown as the conclusion was not affected by the methods chosen. The fractional contribution of organics in the~~
143 ~~signals at $m/z=44$ and $m/z=43$ to total organics (f_{44} and f_{43} , respectively) were also derived~~
144 ~~SOA yields were~~
145 ~~calculated as the ratio of organic aerosol mass formed to the amount of VOC reacted. The concentration of~~
146 ~~organic aerosol was derived using the total aerosol mass concentration measured by SMPS with a density of 1 g cm^{-3}~~
147 ~~(to better compare with previous literature) multiplied by with the mass fraction of organics in total aerosol~~
148 ~~characterized by AMS. The organic aerosol concentration was corrected for the particle wall loss and dilution~~
149 ~~loss using the method described in Zhao et al. (2015). The loss of vapor on the wall was not corrected here.~~

149 SOA yields were calculated as the ratio of organic aerosol mass formed to the amount of VOC reacted.
150 The mass concentration of organic aerosol was derived using the total aerosol volume concentration measured by
151 SMPS multiplied by the volume fraction of organics with a density of 1 g cm^{-3} to better compare with previous
152 literature. In the experiments with added SO_2 , sulfuric acid was formed upon photooxidation. ~~Sulfuric acid was~~
153 ~~and~~ partly neutralized by background ammonia, which was introduced into the chamber mainly due to
154 humidification. The ~~density~~ volume fraction of ~~the aerosol~~ organics was derived ~~using the linear mixing of~~
155 ~~the~~ based on volume additivity using the mass of organics and ammonium sulfate/ammonium bisulfate from
156 AMS and their respective density of organic aerosol (assuming 1.32 g cm^{-3} for organic aerosol from one of our
157 previous studies and the literature (Flores et al., 2014; Eddingsaas et al., 2012a)(Flores et al., 2014)) and the
158 density of ammonium sulfate/ammonium bisulfate (– and the literature (Ng et al., 2007) and $\sim 1.77\text{ g cm}^{-3}$), for
159 ammonium sulfate/ammonium bisulfate). According to the calculations based on the E-AIM model (Clegg et al.,
160 1998; Wexler and Clegg, 2002) (<http://www.aim.env.uea.ac.uk/aim/aim.php>), there ~~were~~ was no aqueous phase
161 formed at the relative humidity in the experiments of this study ~~(– The average RH 28–42% for the period of~~
162 ~~monoterpene photooxidation)–~~ was 28–34% except for one experiment with average RH of 42% RH. The organic
163 aerosol concentration was corrected for the particle wall loss and dilution loss using the method described in
164 Zhao et al. (2015).

165 2.2 Experimental procedure

166 The SOA formation from α -pinene and limonene photooxidation was investigated at different NO_x and
167 SO_2 levels. Four types of experiments were done: with neither NO_x nor SO_2 added (referred to as “low NO_x , low
168 SO_2 ”), with only NO_x added (~ 20 ppb NO , referred to as “high NO_x , low SO_2 ”), with only SO_2 added (~ 15 ppb,
169 referred to as “low NO_x , high SO_2 ”), and with both NO_x and SO_2 added (~ 20 ppb NO and ~ 15 ppb SO_2 , referred
170 to as “high NO_x , high SO_2 ”). For low NO_x conditions, background NO concentrations were around 0.05–0.2 ppb,
171 and NO was mainly from the background photolytic process of Teflon chamber wall (Rohrer et al., 2005). For
172 low SO_2 conditions, background SO_2 concentrations were below the detection limit of the SO_2 analyzer (0.05
173 ppb). In some experiments, a lower level of SO_2 (2 ppb, referred to as “moderate SO_2 ”) was used to test the effect
174 of SO_2 concentration. An overview of the experiments is shown in Table 1.

175 In a typical experiment, the chamber was humidified to $\sim 75\%$ RH first, and then VOC and NO , if
176 applicable, were added to the chamber. Then the roof was opened to start photooxidation. In the experiments with
177 SO_2 , SO_2 was added and the roof was opened to initialize nucleation first and then VOC was added. The particle
178 number concentration caused by SO_2 oxidation typically reached several 10^4 cm^{-3} (see Fig. 2 high SO_2 cases) and

179 after VOC addition, no further nucleation occurred. Adding SO₂ first and initializing nucleation by SO₂
180 photooxidation ensured that enough nucleating particles were present when VOC oxidation started. SO₂
181 concentration decayed slowly in the experiments with SO₂ added and most of the SO₂ was still left (typically
182 around 8 ppb from initial 15 ppb) at the end of an experiment due to its low reactivity with OH. Typical SO₂
183 time series in high SO₂ experiments are shown in Fig S2. The detailed conditions of the experiments are shown in
184 Table S1. The experiments of α -pinene and limonene photooxidation were designed to keep the initial OH
185 reactivity and thus OH loss rate constant so that the OH concentrations of these experiments were more
186 comparable. Therefore, the concentration of limonene was around one-third of the concentration of α -pinene due
187 to the higher OH reactivity of limonene.

188 2.3 Wall loss of organic vapors

189 The loss of organic vapors on chamber walls can influence SOA yield (Kroll et al., 2007; Zhang et al.,
190 2014; Ehn et al., 2014; Sarrafzadeh et al., 2016; McVay et al., 2016; Nah et al., 2016; Matsunaga and Ziemann,
191 2010; Ye et al., 2016; Loza et al., 2010). The wall loss rate of organic vapors in our chamber was estimated by
192 following the decay of organic vapor concentrations after photooxidation was stopped in the experiments with
193 low particle surface area ($\sim 5 \times 10^{-8} \text{ cm}^2 \text{ cm}^{-3}$) and thus low condensational sink on particles. Such method is
194 similar to the method used in previous studies (Ehn et al., 2014; Sarrafzadeh et al., 2016; Krechmer et al., 2016;
195 Zhang et al., 2015). A high-resolution time-of-flight chemical ionization mass spectrometer (HR-ToF-CIMS,
196 Aerodyne Research Inc.) with nitrate ion source ($^{15}\text{NO}_3^-$) was used to measure semi/low-volatile organic vapors.
197 The details of the instrument were described in our previous studies (Ehn et al., 2014; Sarrafzadeh et al., 2016).
198 The decay of vapors started from the time when the roof of the chamber was closed. The data were acquired at a
199 time resolution of 4 s. A typical decay of low-volatile organics is shown in Fig. S3 and the first-order wall loss
200 rate was determined to be around $6 \times 10^{-4} \text{ s}^{-1}$.

201 The SOA yield was not directly corrected for the vapor wall loss, but the influence of vapor wall loss on
202 SOA yield was estimated using the method in the study of Sarrafzadeh et al. (2016) and the details of the method
203 are described therein. Briefly, particle surface and chamber walls competed for the vapor loss (condensation) and
204 the condensation on particles led to particle growth. The fraction of organic vapor loss to particles in the sum of
205 the vapor loss to chamber walls and to particles (F_p) was calculated. The vapor loss to chamber walls was derived
206 using the wall loss rate. The vapor loss to particles was derived using particle surface area concentration,
207 molecular velocity and an accommodation coefficient α_p (Sarrafzadeh et al., 2016). $1/F_p$ (f_{corr}) provides the
208 correction factor to obtain the “real” SOA yield. f_{corr} is a function of particle surface area concentration and
209 accommodation coefficient as shown in Fig. S4. Here a range of 0.1-1 for α_p was used, which is generally in line
210 with the ranges of α_p found by Nah et al. (2016) by fitting a vapor-particle dynamic model to experimental data.
211 At a given α_p , the higher particle surface area, the lower f_{corr} and the lower the influence of vapor wall loss are
212 because most vapors condense on particle surface and vice versa. At a given particle surface area, f_{corr} decreases
213 with α_p because at higher α_p a larger fraction of vapors condenses on particles. An average molecular weight of
214 200 g/mol was used to estimate the influence of vapor wall loss. For the aerosol surface area range in most of the
215 experiments in this study, the influence of vapor wall loss on SOA yield was relatively small ($< \sim 40\%$ for particle

216 surface area larger than $3 \times 10^{-6} \text{ cm}^2 \text{ cm}^{-3}$, Fig. S4). Yet, for the experiments at high NO_x and low SO_2 for α -pinene
217 and limonene, the influence of vapor wall loss on SOA can be high due to the low particle surface area, especially
218 at lower α_p .

219 3 Results and discussion

220 3.1 Chemical scheme: VOC oxidation pathway and RO_2 fate

221 In the photooxidation of VOC, OH and O_3 often co-exist and both contribute to VOC oxidation. because
222 O_3 formation in chamber studies is often unavoidable during photochemical reactions of VOC even in the
223 presence of trace amount of NO_x . In order to study the mechanism of SOA formation, it is helpful to isolate one
224 oxidation pathway from the other. In this study, the reaction rates of OH and ozone with VOC are quantified
225 using measured OH and O_3 concentrations multiplied by rate constants. (time series of VOC, OH, and O_3 are
226 shown in Fig. S5). Typical OH and O_3 concentrations in an experiment were around $(1-15) \times 10^6 \text{ molecules cm}^{-3}$
227 and 0-50 ppb, respectively, depending on the VOC and NO_x concentrations added. For all the experiment in this
228 study, the VOC loss was dominated by OH oxidation over ozonolysis (see Fig. S6 as an example). The relative
229 importance of the reaction of OH and O_3 with monoterpenes in a typical experiment is shown was similar in Fig.
230 S1. The VOC loss was dominated by OH oxidation over ozonolysis. This the low NO_x and high NO_x experiments.
231 At high NO_x , OH was often higher while more O_3 was also produced. The dominant role of OH oxidation in
232 VOC loss makes the chemical scheme simple and it is easier to interpret than cases when both OH oxidation and
233 ozonolysis are important.

234 As mentioned above, RO_2 fate, i.e., the branching of RO_2 loss among different pathways, has an important
235 influence on the product distribution and thus on SOA composition, physicochemical properties, and yields. RO_2
236 can react with NO, HO_2 , RO_2 , or isomerize. The fate of RO_2 mainly depends on the concentrations of NO, HO_2
237 and RO_2 . Here, the loss rates of RO_2 via different pathways were quantified using the measured HO_2 , NO and
238 RO_2 concentrations and the rate constants based on the MCM3.3 (Jenkin et al., 1997; Saunders et al., 2003)
239 (<http://mcm.leeds.ac.uk/MCM>). Measured HO_2 and RO_2 concentrations in a typical experiment are shown in Fig.
240 S2S7 as an example and the relative importance of different RO_2 reaction pathways is compared in Fig. 1, which
241 is similar for both α -pinene and limonene oxidation. In the low NO_x conditions of this study, $\text{RO}_2 + \text{NO}$ dominated
242 the RO_2 loss rate in the beginning of an experiment (Fig. 1a) because a). The trace amount of NO (up to ~ 0.2
243 ppbV) was formed from the photolysis of HONO, which was continuously produced from a photolytic process on
244 the chamber wallwalls throughout an experiment (Rohrer et al., 2005). But later in the experiment, $\text{RO}_2 + \text{HO}_2$
245 contributed a significant fraction (up to $\sim 40\%$) to RO_2 loss because of increasing HO_2 concentration and
246 decreasing NO concentration. In the high NO_x conditions, $\text{RO}_2 + \text{NO}$ overwhelmingly dominated the RO_2 loss rate
247 (Fig. 1b), and with the decrease of NO in an experiment, the total RO_2 loss rate decreased substantially (Fig. 1b).
248 Since the main products of $\text{RO}_2 + \text{HO}_2$ are organic hydroperoxides, more organic hydroperoxides relative to
249 organic nitrate is nitrates are expected in the low NO_x conditions here. The loss rate of $\text{RO}_2 + \text{RO}_2$ was estimated to
250 be $\sim 10^{-4} \text{ s}^{-1}$ using a reaction rate constant of $2.5 \times 10^{-13} \text{ molecules}^{-1} \text{ cm}^3 \text{ s}^{-1}$ (Ziemann and Atkinson, 2012). This
251 contribution is negligible compared to other pathways in this study, although the reaction rate constants of

252 RO₂+RO₂ are highly uncertain and may depend on specific RO₂ (Ziemann and Atkinson, 2012). Note that the
253 RO₂ fate in the low and high NO_x conditions quantified here are further used in the discussion below since the
254 information of RO₂ fate is important for data interpretation of experiments conducted at different NO_x levels
255 (Wennberg, 2013).

256 3.2 Effects of NO_x and SO₂ on new particle formation

257 The effects of NO_x and SO₂ on new particle formation from α -pinene oxidation are shown in Fig. 2a. In
258 low SO₂ conditions, both the total particle number concentration and nucleation rate at high NO_x were lower than
259 those at low NO_x, indicating NO_x suppressed the new particle formation. The suppressing effect of NO_x on new
260 particle formation was in agreement with the findings of Wildt et al. (2014). This suppression is considered to be
261 caused by the increased fraction of RO₂+NO reaction, decreasing the importance of RO₂+RO₂ permutation
262 reactions. RO₂+RO₂ reaction products are believed to be involved in the new particle formation (Wildt et al.,
263 2014; Kirkby et al., 2016) and initial growth of particles by forming higher molecular weight products such as
264 highly oxidized multifunctional molecules (HOM) and their dimers and trimers (Ehn et al., 2014; Kirkby et al.,
265 2016).

266 In high SO₂ conditions, the nucleation rate and total number concentrations were high, regardless of NO_x
267 levels. The high concentration of particles was attributed to the new particle formation induced by H₂SO₄ alone
268 formed by SO₂ oxidation since the new particle formation occurred before VOC addition. The role of H₂SO₄ in
269 new particle formation has been well studied in previous studies (Berndt et al., 2005; Zhang et al., 2012; Sipila et
270 al., 2010; Kirkby et al., 2011; Almeida et al., 2013).

271 Similar suppression of new particle formation by NO_x and enhancement of new particle formation by
272 SO₂ photooxidation were found for limonene oxidation (Fig. 2b).

273 3.3 Effects of NO_x and SO₂ on SOA mass yield

274 3.3.1 Effect of NO_x

275 Figure 3a shows SOA yield at different NO_x for α -pinene oxidation. In order to make different
276 experiments more comparable, the SOA yield is plotted as a function of OH dose instead of reaction time. In low
277 SO₂ conditions, NO_x not only suppressed the new particle formation but also suppressed SOA mass yield.
278 Because NO_x suppressed new particle formation, the suppression ~~of the~~ SOA yield could be attributed to the
279 ~~absence~~lack of ~~nucleation~~new particles as seed and thus the ~~absence~~lack of condensational sink, or to the
280 decrease of condensable organic materials. We further found that when new particle formation was already
281 enhanced by added SO₂, the SOA yield at high NO_x was comparable to that at low NO_x and the difference in
282 SOA yield between high NO_x and low NO_x was much smaller (Fig. 3a). This finding ~~indicates that NO_x can be~~
283 attributed to two possible explanations. Firstly, NO_x did not significantly suppress the formation of low volatile
284 condensable organic materials, although NO_x obviously suppressed the formation of products for nucleation.
285 Secondly, NO_x did suppress the formation of low-volatility condensable organic materials via forming potentially
286 more volatile compounds and in addition to that, the suppressed formation of condensable organic materials was
287 compensated by the presence of SO₂, resulting in comparable SOA yield. Organic nitrates are a group of

288 ~~compounds formed at high NO_x, which have been proposed to be more volatile (Presto et al., 2005; Kroll et al.,~~
289 ~~2006). However, many organic nitrates formed by photooxidation in this study were highly oxidized organic~~
290 ~~molecules (HOMs) containing multi-functional groups besides nitrate group (C₇₋₁₀H₉₋₁₅NO₈₋₁₅). These compounds~~
291 ~~are expected to have low volatility and they are found to have an uptake coefficient on particles of ~1 (Pullinen et~~
292 ~~al., in preparation). Therefore, we conclude that~~ the suppressing effect of NO_x on SOA yield was ~~mainly~~
293 ~~likely~~ due to ~~suppressing~~ nucleation, i.e., ~~to the absence~~ lack of particle surface as condensational sink.

294 For limonene oxidation, similar results of NO_x suppressing the particle mass formation have been found
295 in low SO₂ conditions (Fig. 3b). Yet, in high SO₂ conditions, the SOA yield from limonene oxidation at high NO_x
296 was still significantly lower than that at low NO_x, which is different from the findings for α-pinene SOA. ~~That~~
297 ~~might be caused by either the larger~~ The cause of this difference ~~between the volatility of oxidation is currently~~
298 ~~unknown. Our data of SOA yield suggest that the~~ products formed ~~under different NO_x conditions for~~
299 ~~limonene case compared to α-pinene or by the different ranges of VOC/NO_x for α-pinene (VOC/NO_x=1~~
300 ~~oxidation at high NO_x, see Table 1) and limonene (VOC/NO_x=0.35, seemed to have higher average volatility than~~
301 ~~that at high/low NO_x).~~

302 The suppression of SOA mass formation by NO_x under low SO₂ conditions agrees with previous studies
303 (Eddingsaas et al., 2012a; Wildt et al., 2014; Sarrafzadeh et al., 2016; Hatakeyama et al., 1991). For example, it
304 was found that ~~high~~ concentration of NO_x (tens of ppb) suppressed mass yield of SOA formed from
305 photooxidation of β-pinene, α-pinene and VOC emitted by Mediterranean trees (Wildt et al., 2014; Sarrafzadeh et
306 al., 2016). And on the basis of the results by Eddingsaas et al. (2012a), the SOA yield at high NO_x (referred to as
307 “high NO” by the authors) is lower than at low NO_x in the absence of seed aerosol.

308 Our finding that the difference in SOA yield ~~at~~ between high NO_x ~~is comparable to that at~~ and low NO_x ~~in~~
309 ~~high SO₂ conditions~~ was highly reduced at high SO₂ is also in line with the findings of some previous studies
310 using seed aerosols (Sarrafzadeh et al., 2016; Eddingsaas et al., 2012a). For example, Sarrafzadeh et al. (2016)
311 found that in the presence of seed aerosol, the suppressing effect of NO_x on the SOA yield from β-pinene
312 photooxidation is substantially diminished and SOA yield only ~~varies~~ decreases by 20-30% in the NO_x range of
313 <1 ppb to 86 ppb at constant OH concentrations. The data by Eddingsaas et al. (2012a) also ~~found~~ showed that in
314 presence of seed aerosol, the difference in the SOA yield between low NO_x and high NO_x is much decreased.
315 However, our finding is in contrast with the findings in other studies (~~Presto et al., 2005; Ng et al., 2007; Han et~~
316 ~~al., 2016~~)(Presto et al., 2005; Ng et al., 2007; Han et al., 2016; Stirnweis et al., 2017), who reported much lower
317 SOA yield at high NO_x than at low NO_x in presence of seed. The different findings in these studies from ours
318 may be attributed to the difference in the reaction conditions such as VOC oxidation pathways (OH oxidation vs.
319 ozonolysis), VOC and NO_x concentration ranges, NO/NO₂ ratio as well as OH concentrations as well as organic
320 aerosol loading, which all affect SOA yield. ~~Note that even at “high NO_x” the NO_x concentration in~~ The reaction
321 conditions of this study ~~was much lower than in many previous studies and the NO_x concentration range here was~~
322 ~~more relevant to the anthropogenic-biogenic interactions in the ambient, often differ from those described in the~~
323 literature (see Table S2).

324 The difference in these conditions can result in both different apparent dependence on specific
325 parameters and the varied SOA yield. For example, SOA yield from α-pinene photooxidation at low NO_x in this

326 study appeared to be much lower than that in Eddingsaas et al. (2012a). The difference between the SOA yield in
327 this study and some of previous studies and between the values in the literature can be attributed to several
328 reasons: 1) RO₂ fates may be different. For example, in our study at low NO_x, RO₂+NO account for a large
329 fraction of RO₂ loss while in Eddingsaas et al. (2012a) RO₂+HO₂ is the dominant pathway of RO₂ loss. This
330 difference in RO₂ fates may affect oxidation products distribution. 2) The organic aerosol loading of this study is
331 much lower than that some of previous studies (e.g., Eddingsaas et al. (2012a)) (see Fig. S9). SOA yields in this
332 study were also plotted versus organic aerosol loading to better compare with previous studies (Fig. S8 and S9).
333 3) The total particle surface area in this study may also differ from previous studies, which may influence the
334 apparent SOA yield due to vapor wall loss (the total particle surface area is often not reported in many previous
335 studies to compare with). 4) RH of this study is different from many previous studies, which often used very low
336 RH (<10%). It is important to emphasize that reaction conditions including the NO_x as well as SO₂ concentration
337 range and RH in this study were chosen to be relevant to the anthropogenic-biogenic interactions in the ambient
338 atmosphere. In addition, difference in the organic aerosol density used in yield calculation should be taken into
339 account. In this study, SOA yield was derived using a density of 1 g cm⁻³ to better compare with many previous
340 studies (e.g., (Henry et al., 2012)), while in some other studies SOA yield was derived using different density
341 (e.g., 1.32 g cm⁻³ in Eddingsaas et al. (2012a)).

342 3.3.2 Effect of SO₂

343 For both α -pinene and limonene, SO₂ was found to enhance the SOA mass yield; at given NO_x levels,
344 especially for the high NO_x cases (Fig. 3). The enhancing effect of SO₂ on particle mass formation can be
345 attributed to two reasons. Firstly, SO₂ oxidation induced new particle formation, which provided more surface
346 and volume for further condensation of organic ~~vapor,~~vapors. This is consistent with the finding that the
347 enhancement of SOA yield by SO₂ was more significant at high NO_x when the enhancement in nucleation was
348 also more significant. Secondly, H₂SO₄ formed by photooxidation of SO₂ can enhance SOA formation via acid-
349 catalyzed heterogeneous uptake, an important SOA formation pathway initially found from isoprene
350 photooxidation (Jang et al., 2002; Lin et al., 2012; Surratt et al., 2007) and later also in the photooxidation of
351 other compound such as anthropogenic VOC (Chu et al., 2016; Liu et al., 2016). For the products from
352 monoterpene oxidation, such an acid-catalyzed effect may also occur (Northcross and Jang, 2007; Wang et al.,
353 2012; Lal et al., 2012; Zhang et al., 2006; Ding et al., 2011; Iinuma et al., 2009) and in this study, the particles
354 were acidic with the molar ratio of NH₄⁺ to SO₄²⁻ around 1.5-1.8. ~~As mentioned above, inducing new particle~~
355 ~~formation by SO₂ is especially important at high NO_x conditions, when nucleation, although no aqueous phase~~
356 ~~was suppressed by NO_x. In addition, we formed.~~

357 We found that the SOA yield in the limonene oxidation at a moderate SO₂ level (2 ppb) was comparable
358 to the yield at high SO₂ (15 ppb) when similar particle number concentrations in both cases were formed. Both
359 yields were significantly higher than the yield at low SO₂ (<0.05 ppb, see Fig. ~~S3~~S10). This comparison suggests
360 that the effect in enhancing new particle formation by SO₂ seems to be more important; compared to the particle
361 acidity effect. The role of SO₂ ~~on~~in new particle formation is similar to adding seed aerosol ~~on~~and providing
362 particle surface for organics to condense. Artificially added seed aerosol has been shown to enhance SOA

363 formation from α -pinene and β -pinene oxidation (Ehn et al., 2014; Sarrafzadeh et al., 2016; Eddingsaas et al.,
364 2012a). In some other studies, it was found that the SOA yield from α -pinene oxidation is independent of initial
365 seed surface area (McVay et al., 2016; Nah et al., 2016). The difference in the literature may be due to the range
366 of total surface area of particles, reaction conditions and chamber setup. For example, the peak particle-to-
367 chamber surface ratio for α -pinene photooxidation in this study was 7.7×10^{-5} at high NO_x and low SO_2 , much
368 lower than the aerosol surface area range in the studies by Nah et al. (2016) and McVay et al. (2016). A lower
369 particle-to-chamber surface ratio can lead to a larger fraction of organics lost on chamber walls. Hence, providing
370 additional particle surface by adding seed particles can increase the condensation of organics on particles and
371 thus increase SOA yield. However, once the surface area is high enough to inhibit condensation of vapors on
372 chamber walls, further enhancement of particle surface will not significantly enhance the yield (Sarrafzadeh et al.,
373 2016).

374 ~~The finding that in presence of high SO_2 , the SOA mass yield at NO_x conditions was enhanced to the~~
375 ~~value comparable with that at low NO_x indicates that the suppressing effect of NO_x on SOA mass formation was~~
376 ~~counterbalanced by the presence of SO_2 .~~ Particle acidity may also play a role in affecting the SOA yield in the
377 experiments with high SO_2 . Particle acidity was found to enhance the SOA yield from α -pinene photooxidation at
378 high NO_x (Offenberg et al., 2009) and “high NO” conditions (Eddingsaas et al., 2012a). Yet, in low NO_x
379 condition, particle acidity was reported to have no significant effect on the SOA yield from α -pinene
380 photooxidation (Eddingsaas et al., 2012a; Han et al., 2016). According to these findings, at low NO_x the
381 enhancement of SOA yield in this study is attributed to the effect of facilitating nucleation and providing more
382 particle surface by SO_2 photooxidation. At high NO_x , the effect in enhancing new particle formation by SO_2
383 photooxidation seems to be more important, although the effect of particle acidity resulted from SO_2
384 photooxidation may also play a role.

385 SO_2 has been proposed to also affect gas phase chemistry of organics by changing the HO_2/OH or
386 forming SO_3 (Friedman et al., 2016). In this study, the effect of SO_2 on gas phase chemistry of organics was not
387 significant because of the much lower reactivity of SO_2 with OH compared with α -pinene and limonene
388 (Atkinson et al., 2004, 2006; Atkinson and Arey, 2003) and the low OH concentrations (2-3 orders of magnitude
389 lower than those in the study by Friedman et al. (2016)). Moreover, reactions of RO_2 with SO_2 was also not
390 important because the reaction rate constant is very low ($<10^{-14}$ molecule $^{-1}$ cm 3 s $^{-1}$) (Lightfoot et al., 1992; Berndt
391 et al., 2015). In addition, from the AMS data of SOA formed at high SO_2 no significant organic fragments
392 containing sulfur were found. Also the fragment CH_3SO_2^+ from organic sulfate suggested by Farmer et al. (2010)
393 was not detected in our data. The absence of organic sulfate tracers is likely due to the lack of aqueous phase in
394 aerosol particles in this study. Therefore, the influence of SO_2 on gas phase chemistry of organics and further on
395 SOA yield via affecting gas phase chemistry is not important in this study.

396 The presence of high SO_2 enhanced the SOA mass yield at high NO_x conditions, which was even
397 comparable with the SOA yield at low NO_x for α -pinene oxidation. This finding indicates that the suppressing
398 effect of NO_x on SOA mass formation was compensated to large extent by the presence of SO_2 . This has
399 important implications for SOA formation affected by anthropogenic-biogenic interactions in the real atmosphere
400 as discussed below when SO_2 and NO_x often co-exist in relative high concentrations as discussed below.

401 3.4 Effects of NO_x and SO₂ on SOA chemical composition

402 The effects of NO_x and SO₂ on SOA chemical composition were analyzed on the basis of AMS data. We
403 found that NO_x enhanced nitrate formation. The ratio of the mass of nitrate to organics was higher at high NO_x
404 than at low NO_x regardless of the SO₂ level, and similar trends were found for SOA from α -pinene and limonene
405 oxidation (Fig. 4a). Higher nitrate to organics ratios were observed for SOA from limonene at high NO_x, which is
406 mainly due to the lower VOC/NO_x ratio resulted from the lower concentrations of limonene (7 ppb) compared to
407 α -pinene (20 ppb) (see Table 1). Overall, the mass ratios of nitrate to organics ranged from 0.02 to 0.11
408 considering all the experiments in this study.

409 Nitrate formed can be either inorganic (such as HNO₃ from the reaction of NO₂ with OH) or organic (from
410 the reaction of RO₂ with NO). The ratio of NO₂⁺ ($m/z=46$) to NO⁺ ($m/z=30$) in the mass spectra detected by AMS
411 can be used to differentiate whether nitrate is organic or inorganic (Fry et al., 2009; Rollins et al., 2009; Farmer et
412 al., 2010; Kiendler-Scharr et al., 2016). Organic nitrate was considered to have a NO₂⁺/NO⁺ of ~0.1 and inorganic
413 NH₄NO₃ had a NO₂⁺/NO⁺ of ~0.31 with the instrument used in this study as determined from calibration
414 measurements. In this study, NO₂⁺/NO⁺ ratios ranged from 0.14 to 0.18, closer to the ratio of organic nitrate. The
415 organic nitrate was estimated to account for 57%-77% (molar fraction) of total nitrate considering both the low
416 NO_x and high NO_x conditions. This indicates that nitrate was mostly organic nitrate, even at low NO_x in this
417 study.

418 In order to determine the contribution of organic nitrate to total organics, we estimated the molecular
419 weight of organic nitrates formed by α -pinene and limonene oxidation to be 200-300 g/mol, based on reaction
420 mechanisms ((Eddingsaas et al., 2012b) and MCM v3.3, via website: <http://mcm.leeds.ac.uk/MCM>). We
421 assumed a molecular weight of 200 g/mol in order to make our results comparable to the field studies which used
422 similar molecular weight (Kiendler-Scharr et al., 2016). For this value, the organic nitrate compounds were
423 estimated to account for 7-26% of the total organics mass as measured by AMS in SOA. Organic nitrate fraction
424 in total organics was within the range of values found in a field observation in southeast US (5-12% in summer
425 and 9-25% in winter depending on the molecular weight of organic nitrate) using AMS (Xu et al., 2015b) and
426 particle organic nitrate content derived from the sum of speciated organic nitrates (around 1-17% considering
427 observed variability and 3% and 8% on average in the afternoon and at night, respectively) (Lee et al., 2016).
428 Note that the organic nitrate fraction observed in this study was lower than the mean value (42%) for a number of
429 European observation stations when organic nitrate is mainly formed by the reaction of VOC with NO₃
430 (Kiendler-Scharr et al., 2016).

431 Moreover, we found that the contribution of organic nitrate to total organics (calculated using a
432 molecular weight of 200 g/mol for organic nitrate) was higher at high NO_x (Fig. 4b), although in some
433 experiments the ratios of NO₂⁺ to NO⁺ were too noisy to derive a reliable fraction of organic nitrate. This result is
434 consistent with the reaction scheme that at high NO_x, almost all RO₂ loss was switched to the reaction with NO,
435 which is expected to enhance the organic nitrate formation. Besides organic nitrate, the ratio of nitrogen to carbon
436 atoms (N/C) was also found to be higher at high NO_x (Fig. S4S11). But after considering nitrate functional group
437 separately, N/C ratio was very low, generally <0.01, which indicates majority of the organic nitrogen existed in
438 the form of organic nitrate.

439 The chemical composition of organic components of SOA in terms of H/C and O/C ratios at different
440 NO_x and SO₂ levels was further compared. For SOA from α -pinene photooxidation, in low SO₂ conditions, no
441 significant difference in H/C and O/C was found between SOA formed at low NO_x and at high NO_x within the
442 experimental uncertainties (Fig. 5). The variability of H/C and O/C at high NO_x is large, mainly due to the low
443 particle mass and small particle size. In high SO₂ conditions, SOA formed at high NO_x had the higher O/C and
444 lower H/C, which indicates that SOA components had higher oxidation state. The higher O/C at high NO_x than at
445 low NO_x is partly due to the higher OH dose at high NO_x, although even at same OH dose O/C at high NO_x was
446 still slightly higher than at low NO_x in high SO₂ conditions.

447 For the SOA formed from limonene photooxidation, no significant difference in the H/C and O/C was
448 found between different NO_x and SO₂ conditions (Fig. S5S12), which is partly due to the low signal resulting
449 from low particle mass and small particle size in high NO_x conditions.

450 Due to the high uncertainties for some of the H/C and O/C data, the chemical composition was further
451 analyzed using f_{44} and f_{43} since f_{44} and f_{43} are less noisy (Fig. 6). For both α -pinene and limonene, SOA formed at
452 high NO_x generally has ~~lower~~ lower f_{43} . Because f_{43} generally correlates with H/C in organic aerosol (Ng et al., 2011),
453 lower f_{43} is indicative of lower H/C, which is consistent with the lower H/C at high NO_x observed for SOA from
454 α -pinene oxidation in ~~presence of high concentrations of~~ SO₂ conditions (Fig. 5). The lower f_{43} at high NO_x was
455 evidenced in the oxidation of α -pinene based on the data in a previous study (Chhabra et al., 2011). The lower
456 H/C and f_{43} are likely to be related to the reaction pathways. According to the reaction mechanism mentioned
457 above, at low NO_x a significant fraction of RO₂ reacted with HO₂ forming hydroperoxides, while at high NO_x
458 almost all RO₂ reacted with NO forming organic nitrates. Compared with organic ~~nitrate, hydroperoxides, nitrates,~~
459 hydroperoxides have higher H/C ratio. The same mechanism also caused higher organic nitrate fraction at high
460 NO_x, as discussed above.

461 Detailed mass spectra of SOA were compared, shown in Fig 7. For α -pinene, in high SO₂ conditions,
462 mass spectra of SOA formed at high NO_x generally had higher intensity for CHOgt1 ("gt1" means greater than 1)
463 family ions, such as CO₂⁺ (m/z 44), but lower intensity for CH family ions, such as C₂H₃⁺ (m/z 15), C₃H₃⁺ (m/z 39)
464 (Fig. 7b) than at low NO_x. In low SO₂ conditions, such difference is not apparent (Fig. 7a), partly due to the low
465 signal from AMS for SOA formed at high NO_x as discussed above. For both the high SO₂ and low SO₂ cases,
466 mass spectra of SOA at high NO_x show higher intensity of CHN1 family ions. This is also consistent with the
467 higher N/C ratio shown above. For SOA from limonene oxidation, SOA formed at high NO_x had lower mass
468 fraction at m/z 15 (C₂H₃⁺), 28 (CO⁺), 43 (C₂H₃O⁺), 44 (CO₂⁺), and higher mass fraction at m/z 27 (CHN⁺, C₂H₃⁺),
469 41 (C₃H₅⁺), 55 (C₄H₇⁺), 64 (C₄O⁺) than at low NO_x (Fig. S6S13). It seems that overall mass spectra of the SOA
470 from limonene formed at high NO_x had higher intensity for CH family ions, but lower intensity for CHO1 family
471 ions than at low NO_x. Note that the differences in these m/z were based on the average spectra during the whole
472 reaction period and may not reflect the chemical composition at a certain time.

473 4 Conclusion and implications

474 We investigated the SOA formation from the photooxidation of α -pinene and limonene under different NO_x
475 and SO_2 conditions, when OH oxidation was the dominant oxidation pathway of monoterpenes. The fate of RO_2
476 was regulated by varying NO_x concentrations. We confirmed that NO_x suppressed new particle formation. NO_x
477 also suppressed SOA mass yield in the absence of SO_2 . The suppression of SOA yield by NO_x was mainly likely
478 due to the suppressed new particle formation, i.e., absence of sufficient particle surfaces for organic vapor to
479 condense on at high NO_x . ~~NO_x did not significantly suppress the formation of condensable organics from α -
480 pinene oxidation as evidenced by the similar SOA yield at high and low NO_x in the presence of SO_2 .~~

481 SO_2 enhanced SOA yield from α -pinene and limonene photooxidation. SO_2 oxidation produced high number
482 concentration of particles and compensated for the suppression of SOA yield by NO_x , to a large extent. The
483 enhancement of SOA yield by SO_2 is likely to be mainly caused by facilitating nucleation by H_2SO_4 , although the
484 contribution of acid-catalyzed heterogeneous uptake cannot be excluded.

485 NO_x promoted nitrate formation. The majority (57-77%) of nitrate was organic nitrate at both low NO_x and
486 high NO_x , based on the estimate using the $\text{NO}_2^+/\text{NO}^+$ ratios from AMS data. The significant contribution of
487 organic nitrate to nitrate may have important implications for deriving the hygroscopicity from chemical
488 composition. For example, a number of studies derived the hygroscopicity parameter by linear combination of the
489 hygroscopicity parameters of various components such as sulfate, nitrate, and organics, assuming all nitrates are
490 inorganic nitrate (Wu et al., 2013; Cubison et al., 2008; Yeung et al., 2014; Bhattu and Tripathi, 2015; Jaatinen et
491 al., 2014; Moore et al., 2012; Gysel et al., 2007). Because the hygroscopicity parameter of organic nitrate may be
492 much lower than inorganic nitrate (Suda et al., 2014), such derivation may overestimate hygroscopicity.

493 Organic nitrate compounds are estimated to contribute 7-26% of the total organics using an average
494 molecular weight of 200 g/mol for organic nitrate compounds and a higher contribution of organic nitrate was
495 found at high NO_x . Generally, SOA formed at high NO_x has a lower H/C compared to that at low NO_x . The
496 higher contribution of organic nitrate to total organics and lower H/C at high NO_x than at low NO_x is attributed to
497 the reaction of RO_2 with NO, which produced more organic nitrates relative to organic hydroperoxides formed
498 via the reaction of RO_2 with HO_2 . The different chemical composition of SOA between high and low NO_x
499 conditions may affect the physicochemical properties of SOA such as volatility, hygroscopicity, and optical
500 properties and thus change the impact of SOA on environment and climate.

501 The different effects of NO_x and SO_2 on new particle formation and SOA mass yields have important
502 implications for SOA formation affected by anthropogenic-biogenic interactions in the ambient atmosphere.
503 When an air mass of anthropogenic origin is transported to an area enriched in biogenic VOC emissions or vice
504 versa, anthropogenic-biogenic interactions occur. Such scenarios are common in the ambient atmosphere in many
505 areas. For example, Kiendler-Scharr et al. (2016) shows that the organic nitrate concentrations are high in all the
506 rural sites all over Europe, indicating the important influence of anthropogenic emissions in rural areas which are
507 often enriched in biogenic emissions. ^{14}C analysis in several studies show that modern source carbon, from
508 biogenic emission or biomass burning, account for large fractions of organic aerosol even in urban areas (Szidat
509 et al., 2009; Weber et al., 2007; Sun et al., 2012), indicating the potential interactions of biogenic emissions with
510 anthropogenic emissions in urban areas. In such cases, anthropogenic NO_x alone may suppress the new particle

511 formation and SOA mass from biogenic VOC oxidation, as we found in this study. However, due to the co-
512 existence of NO_x with SO_2 , H_2SO_4 formed by SO_2 oxidation can counteract such suppression of particle mass
513 because regardless of NO_x levels, H_2SO_4 can induce new particle formation especially in the presence of water,
514 ammonia or amine (Berndt et al., 2005; Zhang et al., 2012; Sipila et al., 2010; Almeida et al., 2013; Kirkby et al.,
515 2011; Chen et al., 2012). The overall effects on SOA mass depend on specific NO_x , SO_2 and VOC concentrations
516 and VOC types as well as anthropogenic aerosol concentrations and can be a net suppressing, neutral, or
517 enhancing effect. Such scheme is depicted in Fig. 8. Other anthropogenic emissions, such as primary
518 anthropogenic aerosol and precursors of anthropogenic secondary aerosol, can have similar roles as SO_2 . By
519 affecting the concentrations of SO_2 , NO_x and anthropogenic aerosol, anthropogenic emissions may have
520 important mediating impacts on biogenic SOA formation. Considering the effects of these factors in isolation
521 may cause bias in predicting biogenic SOA concentrations. The combined impacts of SO_2 , NO_x and
522 anthropogenic aerosol are also important to the estimate on how much organic aerosol concentrations will change
523 with the ongoing and future reduction of anthropogenic emissions (Carlton et al., 2010).

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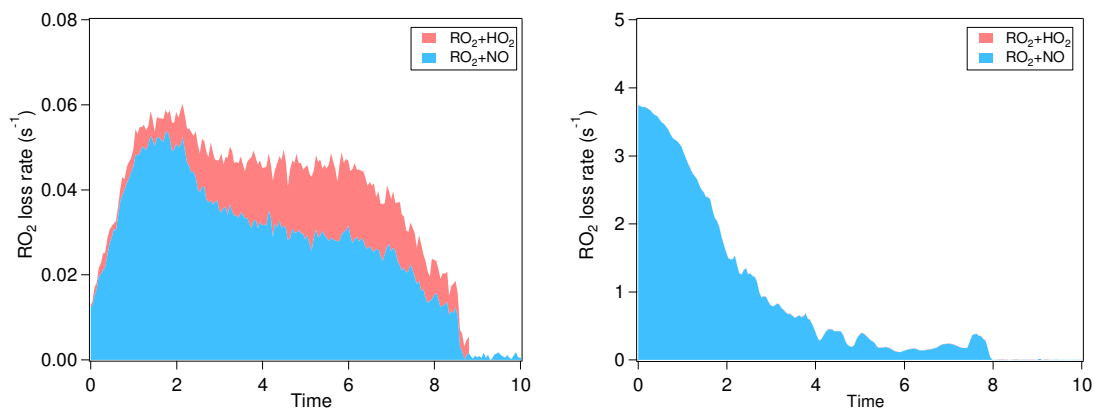
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Table 1 Overview of the experiments in this study

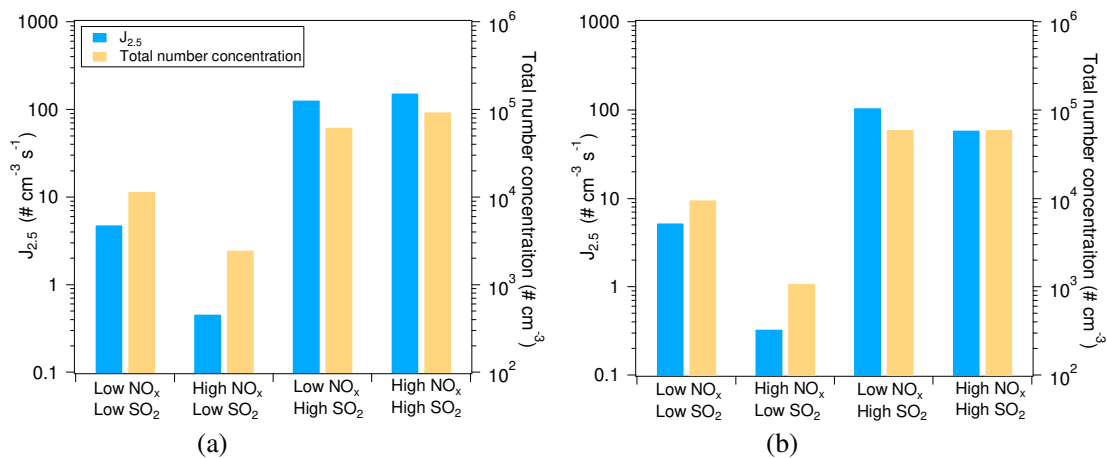
Precursor	SO ₂	NO _x	NO (ppb)	SO ₂ (ppb)
α-pinene	Low SO ₂	Low NO _x	0.05-0.2	<0.05
		High NO _x	~20	<0.05
(~20 ppb)	High SO ₂	Low NO _x	0.05-0.2	~15
		High NO _x	~20	~15
Limonene	Low SO ₂	Low NO _x	0.05-0.2	<0.05
		High NO _x	~20	<0.05
(~7 ppb)	High SO ₂	Low NO _x	0.05-0.2	~15
		High NO _x	~20	~15
	Moderate SO ₂	High NO _x	~20	~2



(a)

(b)

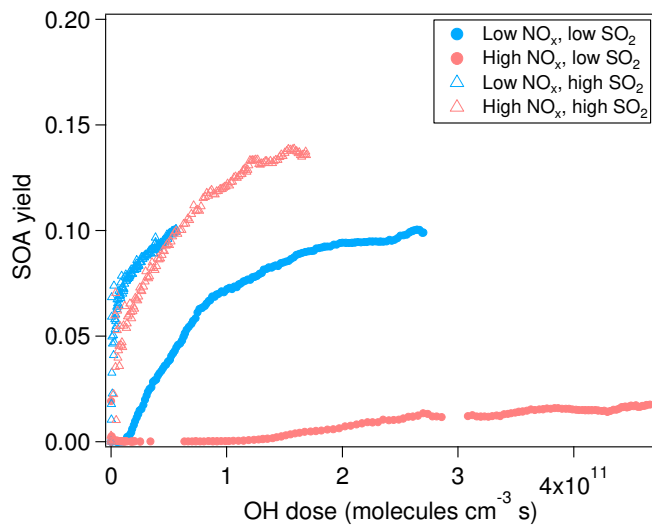
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 928 Figure 1. Typical loss rate of RO₂ by RO₂+NO and RO₂+HO₂ in the low NO_x (a) and the high NO_x (b) conditions
 929 of this study. The experiments at low SO₂ are shown. The RO₂+HO₂ rate is stacked on the RO₂+NO rate. Note
 930 the different scales for RO₂ loss rate in panel a and b. In panel b, the contribution of RO₂+HO₂ is very low and
 931 barely noticeable.



932
 933
 934 Figure 2. Nucleation rates ($J_{2.5}$) and maximum total particle number concentrations under different NO_x and SO_2
 935 conditions for the SOA from α -pinene oxidation (a) and from limonene oxidation (b).

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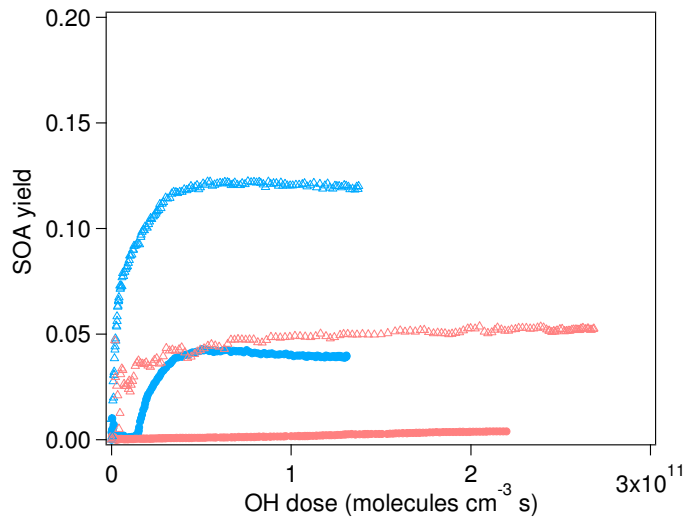
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(a)



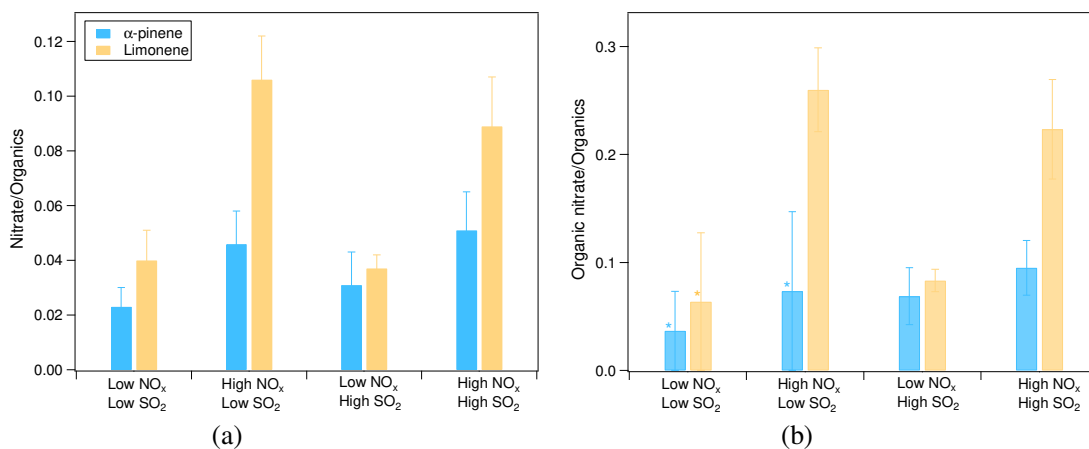
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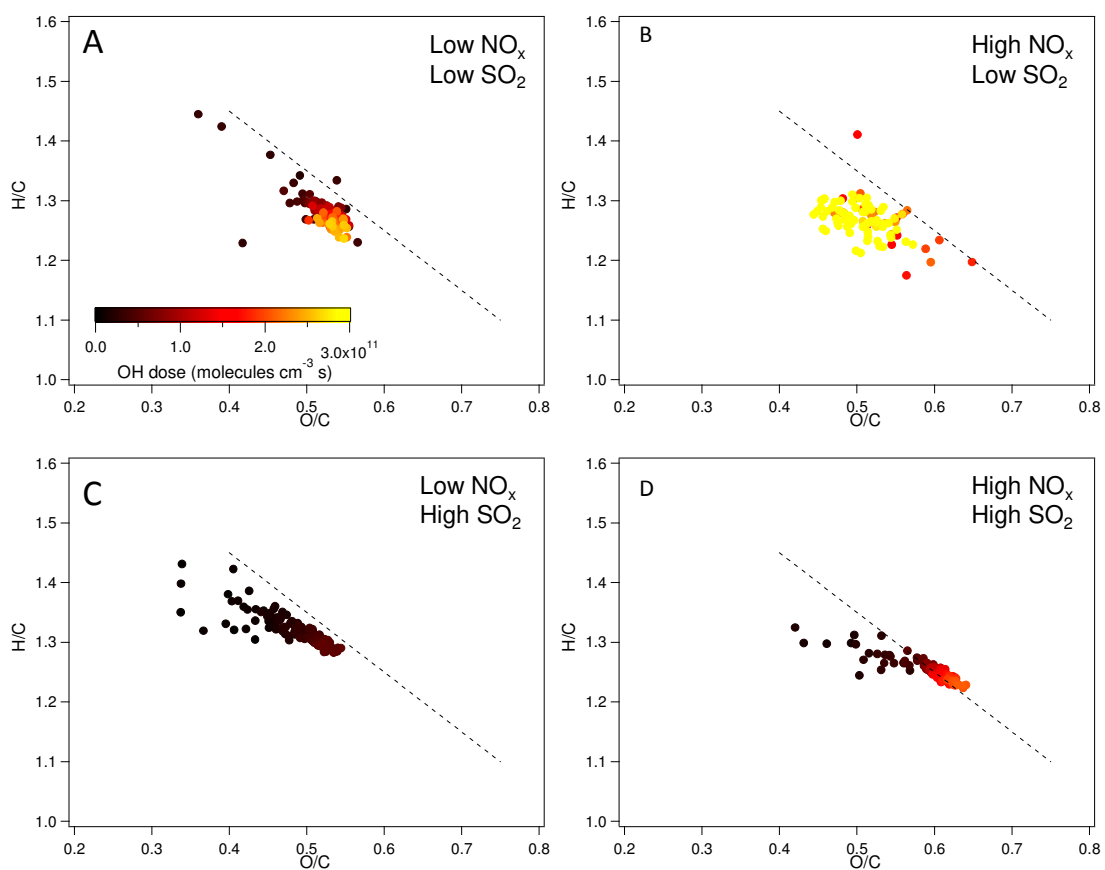
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(b)

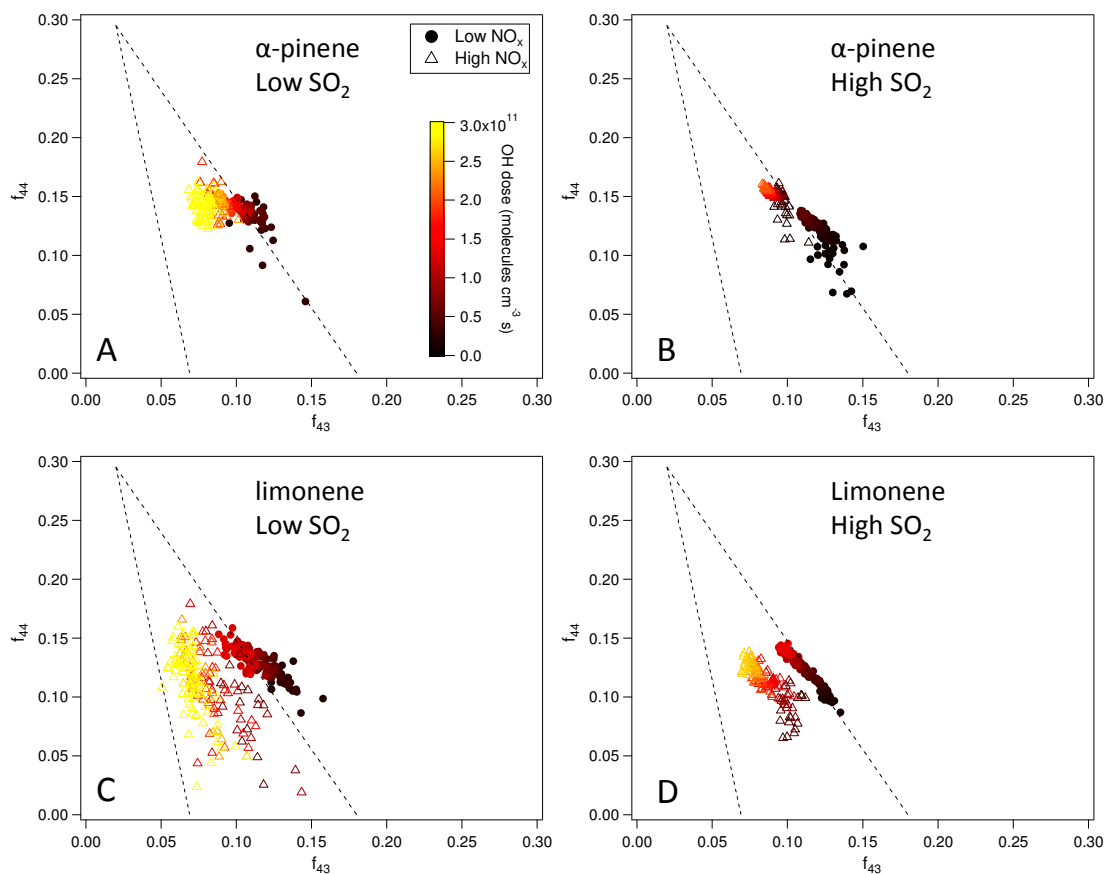
Figure 3. SOA yield of the photooxidation of α -pinene (a) and limonene (b) in different NO_x and SO_2 conditions.



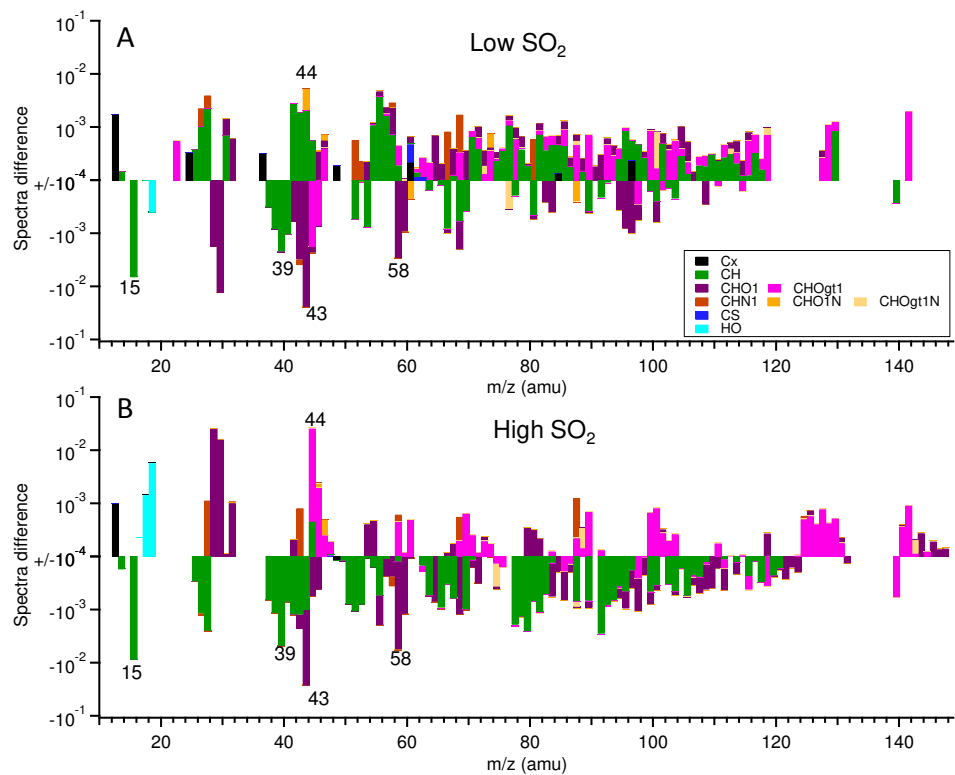
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 946 Figure 4. (a) The ratio of nitrate mass concentration to organics mass in different NO_x and SO_2 conditions; The
 947 average ratios of nitrate to organics during the reaction are shown and error bars indicate the standard deviations.
 948 (b) The fraction of organic nitrate to total organics in different NO_x and SO_2 conditions calculated using a
 949 molecular weight of 200 g/mol for organic nitrate. The average fractions during the reaction are shown and error
 950 bars indicate the standard deviations. In panel b, * indicate the experiments where the ratios of NO_2^+ to NO^+ were
 951 too noisy to derive a reliable fraction of organic nitrate. For these experiments, 50% of total nitrate was assumed
 952 to be organic nitrate and the error bars show the range when 0 to 100% of nitrate are assumed to be organic
 953 nitrate.



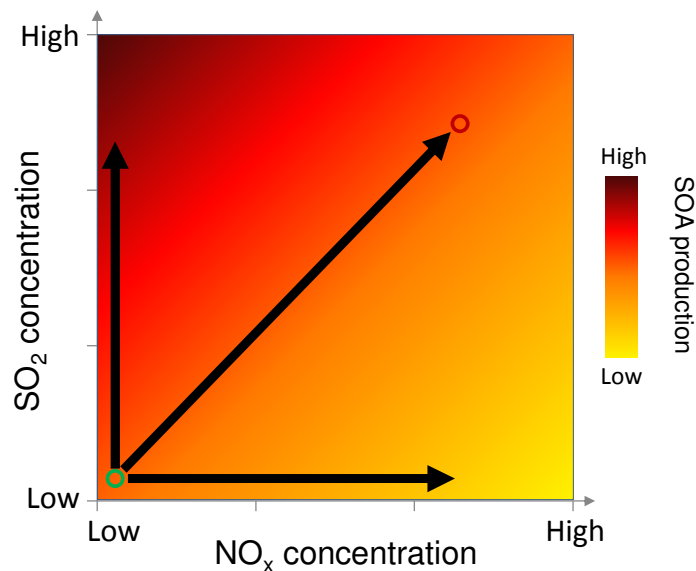
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 956 Figure 5. H/C and O/C ratio of SOA from photooxidation of α -pinene in different NO_x and SO_2 conditions. A:
 957 low NO_x , low SO_2 , B: high NO_x , low SO_2 , C: low NO_x , high SO_2 , D: high NO_x , high SO_2 . The black dashed line
 958 corresponds to the slope of -1.



959
 960 Figure 6. f_{44} and f_{43} of SOA from the photooxidation of α -pinene and limonene in different NO_x and SO_2
 961 conditions. A: α -pinene, low SO_2 , B: α -pinene, high SO_2 , C: limonene, low SO_2 , D: limonene, high SO_2 . Note
 962 that in the low SO_2 , high NO_x condition (panel C), the AMS signal of SOA from limonene oxidation was too low
 963 to derive reliable information due to the low particle mass concentration and small particle size. Therefore, the
 964 data for high NO_x in panel C show an experiment with moderate SO_2 (2 ppb) and high NO_x instead.



965
 966 Figure 7. The difference in the mass spectra of organics of SOA from α -pinene photooxidation between high NO_x
 967 and low NO_x conditions (high NO_x -low NO_x). SOA was formed at low SO_2 (a) and high SO_2 (b). The different
 968 chemical family of high resolution mass peaks are stacked at each unit mass m/z (“gt1” means greater than 1).
 969 The mass spectra were normalized to the total organic signals. Note the log scale of y-axis and only the data with
 970 absolute values large than 10^{-4} are shown.



971
 972 Figure 8. Conceptual schematic showing how NO_x and SO_2 concentrations affect biogenic SOA mass production.
 973 The darker colors indicate higher SOA production. The circle on the bottom left corner indicates biogenic cases
 974 and the circle on the right top corner indicates the anthropogenic cases. And the horizontal and vertical arrows
 975 indicate the effect of NO_x and SO_2 alone. The overall effects on SOA production depend on specific NO_x , SO_2
 976 concentrations and VOC concentrations and speciation.