



- 1 Dependence of the hygroscopicity parameter κ on particle size,
- 2 humidity and solute concentration: implications for laboratory
- 3 experiments, field measurements and model studies
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16 Abstract

17 The hygroscopicity parameter κ has been intensively used in the investigation of 18 the water uptake, cloud condensation nuclei (CCN) activity and chemical 19 composition of atmospheric aerosol particles. A representative value of κ is often 20 assigned to individual species or sources. Such treatment may lead to confusion in 21 closure studies of κ derived from hygroscopic growth factor measurements (κ_{gf}) and 22 CCN activity measurements (κ_{CCN}), and in studies of aerosols at the sub-10 nm size 23 range. Here we show that for particles of the same dry composition, κ may differ as a function of water content, solute concentration and particle size. The concentration-24 25 and size-dependence of κ are demonstrated for representative inorganic and organic 26 compounds, i.e., ammonium sulfate (AS), sodium chloride (NaCl) and sucrose. Our 27 results illustrate that an absolute closure between κ_{gf} and κ_{CCN} should not be expected, 28 and how the deviations observed in field and laboratory experiments can be 29 quantitatively explained and reconciled. The difference between $\kappa_{\rm gf}$ and $\kappa_{\rm CCN}$ 30 increases as particle size decreases reaching up to 40% and 30% for 10 nm AS and 31 NaCl particles, respectively. Moreover, we show that the deviations of κ_{CCN} vary from 32 ~10% for 30 nm and ~40% for 200 nm, indicating a strong dependence on the Köhler 33 models and thermodynamic parameterizations used for instrument calibration (e.g., 34 effective water vapor supersaturation in CCN counter). By taking these factors into 35 account, we can largely explain apparent discrepancies between $\kappa_{\rm gf}$ and $\kappa_{\rm CCN}$ values reported in the scientific literature. Our results help to understand and interpret κ 36 37 values determined at different water vapor ratios and at different size ranges (especially sub-10 nm). We highlight the importance of self-consistent 38 39 thermodynamic parameterizations when using AS for calibration aerosol and taking it 40 as a reference substance representing inorganics in closure study between chemical 41 composition and hygroscopicity of aerosol particles.





42 1. Introduction

43 Hygroscopicity and cloud condensation nuclei (CCN) activity represent the 44 ability of aerosol particles to interact with water, which is essential for the 45 understanding of aerosol climate effects (Andreae and Rosenfeld, 2008). These 46 properties can be described by the Köhler theory (Köhler, 1936), which accounts for 47 both Kelvin and solute effects. Water activity and surface tension are two key 48 parameters in the Köhler equation, both of which are functions of the aerosol 49 composition and solution concentration. The concentration dependence of water activity and surface tension has been determined for many compounds, e.g., 50 51 ammonium sulfate (AS) and sodium chloride (NaCl) etc. (Tang and Munkelwitz, 52 1994; Tang, 1996; Pruppacher and Klett, 1997). Yet we are in lack of such information for a large number of aerosol species and mixtures. 53

54 To describe the relationship between particle dry diameter and CCN activity, 55 Petters and Kreidenweis (2007) proposed a method using a single hygroscopicity 56 parameter κ . The κ parameter has several advantages: (1) representative values of κ 57 may be assigned to a specific aerosol species or source; (2) values of κ for mixtures 58 may be determined from volume-weighted average κ of individual components (Petters and Kreidenweis, 2007); and (3) the experimentally-determined κ has already 59 60 accounted for the impacts of aerosol size, composition and surfactants (Facchini et al., 61 1999). This approach has been proved useful in describing and predicting the CCN activity of single components and aerosol mixtures (Farmer et al., 2015, and 62 references therein), and the simplified parameterization of κ have been implemented 63 in cloud modeling studies (Spracklen et al., 2008; Reutter et al., 2009; Pringle et al., 64 65 2010; Chang et al., 2015).

66 The original purpose of introducing κ is to achieve a simple prediction of critical 67 activation dry diameter and supersaturation for the CCN activation of aerosol





68 particles (Petters and Kreidenweis, 2007). Since representative κ value is assigned to 69 an aerosol species, it is legitimate to ask if the experimentally determined κ can be 70 used to estimate the aerosol composition. Such efforts have been encouraged by the 71 finding of a simple near linear relationship between hygroscopicity parameter κ and 72 the organic mass fraction (Petters and Kreidenweis, 2007; Gunthe et al., 2009), which 73 has been confirmed by results from different locations and sources (Dusek et al., 2010; 74 Cerully et al., 2011; Gunthe et al., 2011; Kawana et al., 2016; Vogel et al., 2016). In 75 addition, the change of κ values was regarded as an indicator for the evolution of 76 chemical composition or mixing state in the aging process of aerosol particles (Chang 77 et al., 2010; Massoli et al., 2010; Wang et al., 2010; Alfarra et al., 2013; Lathem et al., 78 2013; Mei et al., 2013a; Mei et al., 2013b; Zhao et al., 2015), and the size-dependent 79 κ values have also been used as an evidence of size-dependent chemical compositions 80 (Su et al., 2010; Cerully et al., 2011; Rose et al., 2011; Lance et al., 2013).

81 The Köhler equation describes not only the CCN activation under supersaturated 82 conditions but also the hygroscopic growth under subsaturated conditions. 83 Accordingly, κ values can also be determined from hygroscopic growth data 84 measured by hygroscopic tandem differential mobility analyzer (HTDMA). Hence 85 closure studies have been performed to compare κ values determined by hygroscopic 86 growth factor measurements ($\kappa_{\rm sf}$) to those determined by CCN activity measurements 87 (κ_{CCN}). Instead of a closure, most laboratory (Carrico et al., 2008; Massoli et al., 2010; 88 Dusek et al., 2011; Alfarra et al., 2013; Hansen et al., 2015; Dawson et al., 2016) and field measurements (Good et al., 2010; Irwin et al., 2010; Cerully et al., 2011; Wu et 89 90 al., 2013; Bougiatioti et al., 2016; Kim et al., 2016) showed that κ_{CCN} is usually larger 91 than κ_{gf} . Wex et al., (2009) demonstrated that the constant κ value over different 92 concentrations should be reconsidered due to the non-ideality effects in the solution 93 droplet and surface tension variation. Especially in mixed particles, the organic 94 coating or the presence slightly soluble substances can reduce the water transport





95 across the surface by acting as a physical barrier, which may lead to a discrepancy by 96 a factor of 5–10 between $\kappa_{\rm CCN}$ and $\kappa_{\rm gf}$ for secondary organic aerosol (Petters et al., 97 2009; Hansen et al., 2015; Mikhailov et al., 2015; Pajunoja et al., 2015). Hersey et al., 98 (2013) found that the aerosol aging and biomass burning can lead to the discrepancies 99 between $\kappa_{\rm CCN}$ and $\kappa_{\rm gf}$ values that may be related to mixing state. Conversely, higher 100 κ_{gf} compared to κ_{CCN} was also observed, and the discrepancies were more attributed 101 to the instrument technical differences (Irwin et al., 2011). Moreover, recent studies 102 (Keskinen et al., 2013; Kim et al., 2016) tried to related the 103 experimentally-determined κ value to chemical composition during new particle 104 formation and initial growth stage, based on the κ -composition relationship built for 105 much larger particles (dry diameter larger than ~50 nm). Wang et al. (2015), however, 106 reported much smaller κ values for the sub-10 nm aerosols than to the hundred 107 nanometer particles with the same chemical composition. A question is hence raised, 108 if and how accurate a single κ value can be used to represent the whole K öhler curve 109 for the whole dry size range. The answer to this question is especially critical and 110 sensitive for the understanding of multiphase chemistry (Herrmann et al., 2015; 111 Pöschl and Shiraiwa, 2015; Cheng et al., 2016) and for accurate climate modeling 112 (Pajunoja et al., 2015).

113 Another critical issue concerning the application of κ is that its exact value is 114 subject to the selection of thermodynamic parameterizations in K öhler models (Rose 115 et al., 2008; Mikhailov et al., 2013). This is because different thermodynamic 116 parameterizations may lead to different K öhler curves (Cheng et al., 2015; see Fig. 1 117 and references therein), resulting in different calibrations of effective water vapor 118 supersaturation (S_e) in CCN counter (CCNC) and column relative humidity (RH) in 119 HTDMA and consequently different κ values.

120 In this study, we address the question whether the same compound always has 121 the same κ value; if not, what contributes to the difference. In the following, we first





present the theoretical basis and data retrieval methodology. Then we demonstrate the concentration and size dependence of κ for exemplary substances. Finally, we discuss the closure studies between κ_{gf} and κ_{CCN} , and emphasize the importance of the usage of a consistent K öhler model for both calibration with AS and the retrieval of AS κ value when taking it as a reference substance in the inorganic/organic mixing ratio analyses.

128 2. Methodology

129 2.1 к-К öhler model

The K öhler theory describes the equilibrium saturation ratio (*s*) over a spherical
aqueous droplet (K öhler, 1936):

132
$$s = a_{\rm w} \exp(\frac{4\sigma_{\rm sol}v_{\rm w}}{RTD_{\rm d}gf}) \tag{1}$$

where a_w is the water activity, σ_{sol} is the surface tension of solution droplet, v_w is the partial molar volume of water, D_d is the particle dry diameter (mass equivalent diameter), *gf* is the growth factor of particle diameter relative to the dry particle diameter (*gf* = D/D_d), which is related with solute concentration by concentration-dependent solution density. *R* and *T* are the universal gas constant and absolute temperature, respectively.

For AS and NaCl aerosols, we choose the original models used by Biskos et al. (2006a; 2006b), which agree well with the observation data from HTDMA experiments (Cheng et al., 2015). a_w and σ_{sol} are expressed as a function of solute mass fraction x_{s} , which are derived from the best fit to the literature measurement data (Tang and Munkelwitz, 1994; Tang, 1996; Pruppacher and Klett, 1997). The parameterizations are detailed in the Appendix.





145 Petters and Kreidenweis (2007) proposed a hygroscopicity parameter κ for the 146 parameterization of water activity:

147
$$\frac{1}{a_{\rm w}} = 1 + \kappa \frac{V_{\rm s}}{V_{\rm w}}$$
(2)

148 where $V_{\rm s}$ and $V_{\rm w}$ are the volumes of the dry aerosol particle and water, respectively. 149 For ideal solution, κ have a unique value for certain species while it may vary for 150 non-ideal solution.

151 Substituting the expression of κ into Eq. (1), we have the κ -K öhler equation as:

152
$$s = \frac{gf^3 - 1}{gf^3 - (1 - \kappa)} \exp(\frac{4\sigma_{\rm sol}v_{\rm w}}{RTD_{\rm d}gf})$$
(3)

153 Due to the lack of thermodynamic data for the mixed or unknown system, 154 surface tension of water (σ_w) is used instead of σ_{sol} and v_w is simplified as M_w/ρ_w in 155 the κ calculation (Petters and Kreidenweis, 2007; Su et al., 2010), where M_w and ρ_w 156 are the molar mass and density of water, respectively.

157 **2.2 Determination of** *κ*

To elucidate the concentration and size dependence of κ , we use AS, NaCl and sucrose as exemplary substances. The 'real' *s-gf* relations for these substances at different sizes (i.e., D_d) are first calculated by Eq. (1) with well-documented thermodynamic data (i.e., a_w and σ_{sol}). We then retrieve the corresponding κ values for each s-*gf*- D_d pair with Eq. (4):

163
$$\kappa = \frac{gf^3 - 1}{s} \exp(\frac{A}{D_d gf}) - gf^3 + 1 \text{ and } A = \frac{4\sigma_w M_w}{\rho_w RT}$$
(4)

164 **3. Results and discussion**

165 **3.1 Concentration and size dependence of** κ





167 In reality, the aerosol or cloud droplets are often not ideal solutions, especially at concentrated state, and the Zdanovskii-Stokes-Robinson (ZSR) volume additivity 168 169 assumption would not hold anymore. To maintain the κ -K öhler equation (Eq.3), these 170 non-ideality effects have to be compensated, which is reflected by a change in κ . That 171 is why a single species does not correspond to a unique κ value. In addition, since $\sigma_{\rm w}$ 172 is often used instead of σ_{sol} , their difference also needs to be compensated through a 173 change in κ values. In a word, the κ in practical application is subject to both the 174 non-ideality effect and simple treatment of surface tensions. The nano-size effects 175 (e.g., the Tolman effect) on the thermodynamic properties of aerosol particles may 176 result in additional size dependence as well (Cheng et al., 2015).

177 **Concentration dependence.** Figure 1 shows the κ values as a function of solute 178 concentration (expressed as gf or solute molality μ_s) for different aerosol particles with dry diameters of 10 nm, 50 nm and 100 nm, respectively. Here the change of gf 179 180 reflects the change of solution concentration, i.e., larger gf corresponds to more dilute 181 solution (see Eq. A5). Our theoretical calculations illustrate a strong concentration 182 dependence of κ values for AS, NaCl and sucrose. Similar findings in sub-saturation 183 range (RH < 100%) have been shown by Mikhailov et al. (2009; 2013), as water 184 activity dependent κ for AS and levoglucosan. However, in some cases, κ does not show strong concentration dependence, e.g., oxalic acid (Mikhailov et al., 2009) and a 185 186 mixture system of AS and five dicarboxylic acids (Marcolli et al., 2004), which may 187 infer ideal solution or result from the compensation of simplified surface tension and 188 solution non-ideality. For AS and NaCl aerosols (Fig. 1a-b), κ first decreases with increasing gf (decreasing μ_s). After reaching a minimum, κ gradually increases and 189 190 finally approaches a plateau. For example, the κ values of 100 nm AS particles vary 191 from 0.57 at the deliquescence point (RH~80%) to a minimum value of ~0.42, and 192 retain ~0.48 at larger gf. The concentration dependence of NaCl shows a similar 193 shape as AS with κ values eventually approaching ~1.4 (Petters and Kreidenweis,





- 194 2007). However, κ of sucrose aerosols has a different concentration dependence
- 195 which shows a monotonic decrease down to ~0.08. Compared to the 100 nm aerosols,
- 196 the 10 nm and 50 nm aerosols show similar concentration dependence but lower κ 197 values.
- 198 Size dependence. Beside the concentration dependence, Fig. 1 also shows a size 199 dependence of κ . This feature can be more clearly seen in Fig. 2, where κ values are 200 plotted against the particle diameter. The same color-coding represents the same 201 solution concentration, which helps to separate the size effect from the concentration 202 effect. Given the same gf, the κ values increase with increasing particle diameter and 203 finally approaching a plateau, i.e., ~0.56 for AS (gf~1.5), ~1.8 for NaCl (gf~2.0) and 204 ~0.14 for sucrose (gf ~1.2), respectively. The size dependence of κ is more prominent 205 for AS aerosols, especially for smaller particles. For the same concentration (gf of \sim 206 1.5), the κ values of AS vary from 0.40 for 6 nm to 0.51 for 20 nm, and then slowly 207 grow to 0.56 for 300 nm.

208 3.2 Comparison of HTDMA and CCNC measurements

209 3.2.1 Theoretical calculation

According to Eq. (4), κ can be derived experimentally from both hygroscopic growth and CCN activity measurements. Thus direct comparison of κ_{gf} and κ_{CCN} were often performed. The concentration and size dependence of κ , however, challenges the possibility of reaching a closure of κ value determined from different RH or saturation conditions.

215 On the basis of reliable thermodynamic models (see Appendix), we calculate the 216 theoretical κ values (AS, NaCl and sucrose) at RH=90% as for hygroscopic growth 217 factor measurement and at empirical critical supersaturation (S_{cri}) as for CCN activity 218 measurement. Indeed, the retrieved κ is a combined presence of both concentration 219 and size dependence. As demonstrated in Fig. 3 (left panel), κ_{gf} decrease with





220 increasing particle dry diameter while $\kappa_{\rm CCN}$ shows a contrary trend. The gap between

221 HTDMA and CCNC results narrows down with the increasing particle dry diameter

which is also reflected in Fig. 1. Finally, both of the predicted κ approach a value of

~0.48 and ~1.4 for AS and NaCl particles, respectively, while no convergence is
observed for sucrose at 200 nm.

225 We also notice weaker size dependence toward the larger size range. To identify 226 the size range with negligible size effects, we calculate the size dependence of κ 227 values as shown in the right panel of Fig. 3. The size dependence is more prominent 228 for smaller aerosols. For example, a relative variation $(dlog \kappa/dlog D_d)$ of ~0.24 is 229 found for 10 nm AS aerosols while a $dlog \kappa/dlog D_d$ for 50 nm aerosols is only 0.04 230 (CCN activity methods). This size effect should be specifically taking into account 231 when study the properties of nanoparticles during new particle formation and initial 232 growth.

233 3.2.2 Revisit of literature data

234 Figure 4 summarized the comparisons of κ values from HTDMA and CCNC 235 measurements of field and laboratory aerosols (~50-300 nm) in literatures. Different 236 from our theoretical calculations ($\kappa_{\rm CCN} < \kappa_{\rm gf}$), the literature comparison generally 237 shows larger $\kappa_{\rm CCN}$ than $\kappa_{\rm gf}$. Though it was mostly attributed to potential measurement uncertainty (Good et al., 2010) and non-ideality of solution (Wex et al., 2009), we 238 239 provide an alternative explanation that the selection of parameterization methods in 240 the instrument calibration (specifically for CCN activation experiments) may 241 contribute largely to such differences.

Rose et al. (2008) demonstrated the commonly used Köhler models and thermodynamic parameterizations for the calibration of the S_e in CCNC. The results showed that the relative deviations of theoretical S_e are as high as 25% and 12% for AS and NaCl particles, which is mainly caused by the different parameterizations for a_w in dilute aqueous solutions of these two salts. Figure 5a-b shows the deviations of





activity parameterizations (AP) models for AS particles, such as Aerosol Inorganics 247 Model (AIM)-based (referred as AP3) and Electrodynamic Balance method 248 249 (EDB)-based (referred as AP1, roughly corresponding to the models used in this study, see details in the Appendix) a_w parameterizations (Rose et al., 2008). For the range of 250 solute molality up to 0.3 mol kg⁻¹, the a_w differences are on the order of ~10⁻³, 251 resulting in the relative differences of retrieved Se vary between 8% and 18% in the 252 253 size range of 20-200 nm (Figs. 5c and 5e), the size range where the CCNC calibration 254 commonly performed (Rose et al., 2008). Accordingly, the estimated hygroscopicity κ 255 is higher when applying AP3 model in CCN activity measurements compared to AP1 256 model ($\kappa_{AS} \approx 0.66$ with AP3 and $\kappa_{AS} \approx 0.48$ with AP1 at 200 nm). As demonstrated in 257 Figs. 3a and 6a, the deviations appear to be size dependence, varying from $\sim 10\%$ for 258 30 nm to ~40% for 200 nm. Conversely, the influences of selecting different AP 259 models will largely cancel out when performing RH calibration in HTDMA studies 260 (Fig. 5d-e). In other words, the predicted κ values are not sensitive to the selected 261 thermodynamic parameterizations and Köhler models in hygroscopic growth 262 measurements.

263 Although previous studies attributed the inconsistencies between κ_{ef} and κ_{CCN} to 264 the non-ideality effects in the solution droplet and surface tension variation, etc. (Wex et al., 2009; Wu et al., 2013; Hong et al., 2014; Pajunoja et al., 2015; Zhao et al., 265 2016), we find that the key factor to explain the differences is indeed the distinct 266 267 Köhler models and thermodynamic parameterizations used in the CCNC calibration. AIM-based Köhler model (AP3) are recommended because it provides an accurate 268 prediction of a_w for highly dilute solution (Rose et al., 2008). However, our 269 270 theoretical calculations indicate that the AP3 model could lead to a systematic 271 overestimation of up to ~40% in predicting $\kappa_{\rm CCN}$ than $\kappa_{\rm gf}$. In addition, for certain 272 chemical compounds, their hygroscopic behavior in subsaturated and supersaturated 273 conditions is innate distinct, with no apparent closure between κ_{gf} and κ_{CCN} . For





274 example, we find that the κ_{gf} exceeding κ_{CCN} by a factor of 1.5 for sucrose at particle

size larger than 50 nm (Fig. 3e).

276 Our results demonstrate that besides the measurement uncertainty, the 277 discrepancies in previous hygroscopicity closure studies may be to a large extent reconciled by the systematic calibration differences and the inherent concentration 278 279 dependence of κ for specific compounds. Based on our calculations, the ratios of κ_{CCN} 280 to $\kappa_{\rm gf} (\kappa_{\rm CCN}/\kappa_{\rm gf})$ may vary from 1.1 to 1.4 due to systematic calibration differences 281 (Fig. 6a) and from 0.7 to 1.4 due to the inherent concentration dependence for 282 different chemical composition (Figs. 3a and 3e). Taking these effects into account, 283 we find that the observed discrepancy between $\kappa_{\rm gf}$ and $\kappa_{\rm CCN}$ could be partially 284 explained (shaded area in Fig. 6b). Therefore, we suggest that hygroscopicity studies 285 should always report exactly which Köhler models and thermodynamic 286 parameterizations to ensure the comparability of results.

287 3.3 Application to atmospheric closure study between chemical 288 composition and hygroscopicity of aerosol particles

In practice, with precise measurement techniques and reliable calibration procedure, we can accurately measure the particle *gf* at certain RH (HTDMA) and determine the D_{d} - S_{e} relation (CCNC). Hence, the predicted κ value has already considered the impacts of aerosol size, composition and unknown thermodynamic properties, and it can well describe the hygroscopicity and CCN activity. Still, the following issues should be addressed when applying κ :

1. The simple near linear κ -composition relationship establishes the fundamental for estimating κ or mass fraction of organic or inorganic compounds (Dusek et al., 2010; Mei et al., 2013a; Mei et al., 2013b; Wang et al., 2015; Kim et al., 2016; Vogel et al., 2016). In view of such application, the self-consistent approaches should be strictly followed to determine κ value, especially for AS, which is the commonly





300 adopted calibration aerosol and the dominating inorganics of ambient aerosol 301 particles. As previously demonstrated, the retrieval of S_{cri} is very sensitive to the 302 selection of Köhler models and thermodynamic parameterizations, resulting in 303 the representative κ for AS particles vary in a wide range (~0.4-0.8) in closure study 304 (Petters and Kreidenweis, 2007; Rose et al., 2008; Mikhailov et al., 2009; Mikhailov 305 et al., 2013; Wang et al., 2015). Therefore, the chosen κ value should always be 306 correspondingly associated with the calibration method. Otherwise, the inconsistent 307 thermodynamic data will introduce additional uncertainty in retrieving κ value, which 308 may in turn lead to the unreliable estimation of organics/inorganics fractions or 309 hygroscopicity κ for unknown chemical component. For instance, the bias in the 310 estimation of organic mass fraction is within 10% when choosing the identical 311 thermodynamic data to calibrate nano-CCNC and determine κ value for 2.5 nm AS 312 particles, as shown in Fig.9 of Wang et al. (2015). This uncertainty can reach a level 313 of 50% in case of inconsistent parameters are used.

314 2. Size-resolved aerosol hygroscopicity and its link to chemical composition 315 were also investigated in many closure studies (Dusek et al., 2010; Bezantakos et al., 316 2013; Wu et al., 2013; Liu et al., 2014; Zhang et al., 2014). However, only one κ 317 value is commonly used to describe hygroscopicity of the typical inorganics, i.e., ~ 0.6 318 for AS and ~0.7 for NH₄NO₃ in the entire investigated size range. According to our 319 simulation, this hypothesis may not hold for smaller aerosols due to the significant 320 size dependence of κ at that size range (Fig. 3). Experimentally determined κ value is found to be 0.43 for 10 nm AS particles, which is approximately 11% lower than that 321 322 for particles larger than 100 nm. This gap becomes more obvious for sub-10 nm 323 aerosol particles, and the relative deviation is up to $\sim 45\%$. These findings suggest 324 when performing hygroscopicity closure studies, special attention should be paid to 325 size-resolved κ values used in the calculation to obtain accurate results, especially for 326 aerosol particles in nano-size range. In addition, there exist large discrepancies





between observations and model predictions of the hygroscopic growth of salt
nanoparticles when ignoring particle size effects in K öhler models (Biskos et al.,
2006b; Cheng et al., 2015). This prompts the precise thermodynamic data for
nanoparticles needs to be investigated more systematically.

331 4. Conclusion

In this study, we present the differences in particle hygroscopicity behavior (represented as κ) in subsaturated and supersaturated conditions. The proof-of-principle demonstration is performed with AS, NaCl and sucrose particles. Our results indicate that the theoretical κ value shows relatively strong concentration and size dependence, suggesting its value for individual chemical substance should be specifically assessed according to the concerned particle size and saturation regime, especially for sub-10 nm particles.

339 We demonstrate that the observed κ is a compromised product between its 340 inherent concentration dependence (i.e., solution non-ideality & simplification of 341 surface tension) and systematic calibration differences (i.e., different thermodynamic 342 parameterizations in Köhler models). By taking these factors into account, we can 343 largely explain the observed inconsistency of κ values derived from hygroscopic 344 growth factor and CCN activity measurements in literatures. Our results highlight the 345 importance of self-consistent thermodynamic data in the closure studies concerning particle hygroscopicity. 346





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351 Appendix

- 352 In this appendix, we present the parameterizations to retrieve K öhler curve of AS,
- 353 NaCl and sucrose aerosol particles.
- 354 1. AS and NaCl
- 355 Water activity a_w was derived from the single particle experiments by the
- Electrodynamic Balance (EDB) method (Tang and Munkelwitz, 1994; Tang, 1996),
- 357 which can be expressed as polynomial fit functions of solute mass fraction x_s :

358
$$a_{w} = 1 + \sum_{q} A_{q} (100 \cdot x_{s})^{q}$$
(A1)

359 The polynomial coefficients A_q for AS at 298 K are $A_1 = -2.715 \times 10^{-3}$, $A_2 = 3.113 \times 10^{-5}$,

- 360 $A_3 = -2.336 \times 10^{-6}$, and $A_4 = 1.412 \times 10^{-8}$. The polynomial coefficients A_q for NaCl at
- 361 298 K are $A_1 = -6.366 \times 10^{-3}$, $A_2 = 8.624 \times 10^{-5}$, $A_3 = -1.158 \times 10^{-5}$, and $A_4 = 1.518 \times 10^{-7}$.

362 For the surface tension (σ_{sol}) of aqueous AS and NaCl droplet, Pruppacher and

363 Klett (1997) proposed the following parameterization:

364
$$\sigma_{\rm sol} = 0.072 + \frac{2.34 \times 10^{-2} \cdot x_s}{1 - x_s} \text{ for AS}$$
(A2)

365
$$\sigma_{\rm sol} = 0.072 + \frac{2.9 \times 10^{-2} \cdot x_s}{1 - x_s} \quad \text{for NaCl} \tag{A3}$$

of which the validated range is $x_s < 0.78$ for AS and $x_s < 0.45$ for NaCl.

367 Solution density ρ_{sol} is also a function of x_s , which is parameterized based on the 368 EDB measurements (Tang and Munkelwitz, 1994; Tang, 1996):

369
$$\rho_{\rm sol} = \rho_{\rm w} + \sum_{q} d_q (100 \cdot x_s)^q \tag{A4}$$

The polynomial coefficients d_q for AS at 298 K are $d_1 = 5.92$, $d_2 = -5.036 \times 10^{-3}$ and d_3 371 = 1.024×10^{-5} . The polynomial coefficients d_q for NaCl at 298 K are $d_1 = 7.41$, $d_2 = -3.741 \times 10^{-2}$, $d_3 = -2.252 \times 10^{-3}$, and $d_4 = -2.06 \times 10^{-5}$.





373 Growth factor *gf* can be calculated as:

374
$$gf = \left(\frac{\rho_s}{x_s \rho_{sol}}\right)^{1/3} \tag{A5}$$

375 where ρ_s is the density of dry solute.

To be consistent with previous work (Biskos et al., 2006a), the molar volume of pure water was used instead of the partial molar volume and shaper factor correction for dry equivalent spherical mobility diameter were performed based on the methods of Biskos et al. (2006a) for AS and DeCarlo et al. (2004) and Biskos et al. (2006b) for NaCl, respectively.

We also compare the AP1 and AP3 models (section 3.2.2) by following the parameterizations described in Rose et al. (2008). Note that the difference between our present model and AP1 is the different selection of σ_{sol} parameterization. We adopt the parameterization proposed by Pruppacher and Klett (1997) due to the better agreement between calculations and observations for nanoparticles down to 6 nm (Cheng et al., 2015).

387 2. Sucrose

388 For sucrose aerosols, water activity a_w was calculate based on the model provided 389 by Chen (1989):

390
$$a_{\rm w} = \frac{1}{1 + 0.018 \times 0.1136 \times \mu_s^{0.955} \mu_s} \tag{A6}$$

391 where μ_s is the solute molality.

392 Under the assumption of volume additivity (Mikhailov et al., 2004), ρ_{sol} can be 393 calculated by:

394
$$\rho_{\rm sol} = \left(\frac{1-x_s}{\rho_{\rm w}} + \frac{x_s}{\rho_s}\right)^{-1} \tag{A7}$$

395 where $\rho_{\rm w}$ is the density of pure water.

396 Surface tension (σ_{sol}) parameters were adopted by the work of Matubayasi and





- 397 Nishiyama (2006), which measured the surface tension (σ_{sol}) of aqueous sucrose
- 398 solation in the concentration range 0 to 0.5 mol kg^{-1} using drop volume method.
- 399 Growth factor *gf* can be calculated with Eq. A5.





400 Figures



402 Figure 1: Concentration dependence of the hygroscopicity parameter κ calculated for 403 deliquesced aerosol particles of (a) ammonium sulfate (AS), (b) sodium chloride 404 (NaCl) and (c) sucrose with dry diameters of 10 nm (red), 50 nm (blue) and 100 nm 405 (green). Lower limit values of the particle diameter growth factors (gf) correspond to 406 the deliquescence point of crystalline particles. Open circles and triangles indicate the values of κ and solute molality (or diameter gf) calculated at 90% RH and at the 407 408 theoretical critical supersaturation of CCN activation (S_{cri}) , respectively. 409 Thermodynamic parameterizations are described in the Appendix.







411 **Figure 2**: Size dependence of the hygroscopicity parameter κ calculated as a function 412 of particle dry diameter (D_d , x axis) and growth factor (gf, color coding) for 413 deliquesced (a) ammonium sulfate (AS), (b) sodium chloride (NaCl) and (c) sucrose 414 aerosols. The pink lines represent the dependence of κ on D_d for fixed gf values: (a) gf415 = 1.5 for AS, (b) gf = 2.0 for NaCl, and (c) gf = 1.2 for sucrose.







417 Figure 3: Left Panel: Size dependence of the hygroscopicity parameter κ calculated as a function of dry diameter (D_d, x axis) for different types of aerosol particles at a 418 419 typical level of relative humidity applied in HTDMA measurements (RH = 90%, solid 420 lines) and at the critical supersaturation of CCN activation (S_{cri} , dashed and 421 dash-dotted lines): (a) ammonium sulfate (AS), (c) sodium chloride (NaCl) and (e) 422 sucrose. Right panel: The relative sensitivity of κ to changes in dry particle diameter 423 $(dlog \kappa/dlog D_d)$. Solid lines and dash lines were determined with the parameterizations described in the Appendix. Pink dash-dotted lines are determined with the Aerosol 424 425 Inorganics Model (AIM)-based parameterizations (Clegg et al., 1998). Gray dash 426 lines are to guide the eye.







428 **Figure 4**: Comparison of reported κ values from field measurements (circles and 429 solid lines) and laboratory experiments (squares) based on hygroscopic growth factor 430 (x axis) and CCN activity measurements (y axis). The gray dashed line shows 1:1 431 ratio. The size of symbols indicates the particle diameter. Note for the studies which 432 did not provide the specific particle size (Carrico et al., 2008; Hersey et al., 2013; 433 Zhao et al., 2016), we use 100 nm to present the results.







435 Figure 5: Left panel: (a-b) Dependence of water activity (a_w) on solute concentration 436 (molality μ_s) in aqueous ammonium sulfate (bulk solution). Pink triangles indicate the 437 electrodynamic balance (EDB) measurement data (Clegg et al., 1995). Blue and green lines represent different parameterizations of a_w as a function of μ_s based on EDB 438 439 method (AP1) and on the Aerosol Inorganics Model (AP3). Right panel: (b) 440 Comparison of results when applying different AP models in calibrating (c) the 441 effective supersaturation (Se) in a CCNC instrument and (d) the effective relative 442 humidity (RH) in a HTDMA instrument. (e) Ratio of calibration values (AP1 to AP3 models) for HTDMA (e.g., RH at gf=1.7) and CCNC (S_e). The corresponding 443 concentration and saturation ranges for HTDMA and CCNC measurements are 444 445 respectively marked as yellow and pink.







447 Figure 6: (a) Comparison of retrieved hygroscopicity κ using different Activity 448 Parameterization (AP) model (Rose et al., 2008) for the calibration of HTDMA and 449 CCNC. The dashed line shows the 1:1 ratio, and the solid lines correspond to 10% 450 and 50% uncertainties. (b) The shaded area represents the region where the measured 451 κ values might be located when considering systematic calibration differences 452 ($\kappa_{\rm CCN}/\kappa_{\rm gf} \sim 1.1-1.5$) and inherent concentration dependence ($\kappa_{\rm CCN}/\kappa_{\rm gf} \sim 0.7-1.4$) of κ . 453 The upper and lower bounds corresponding to the fit lines of $\kappa_{CCN}=2.1 \kappa_{gf}$ and 454 $\kappa_{\rm CCN}$ =0.7 $\kappa_{\rm gf}$, respectively. Reported κ values from field measurements and laboratory 455 experiments corresponding to Fig. 4.





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