

A new balance formula to estimate new particle formation rate: reevaluating the effect of coagulation scavenging

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Abstract. A new balance formula to estimate new particle formation rate is proposed. It is derived from aerosol general dynamic equation in the discrete form and then converted into an approximately continuous form for analysing data from new particle formation (NPF) field campaigns. The new formula corrects the underestimation of the coagulation scavenging effect occurred in previously used formulae. It also clarifies the criteria in determining upper size bound in measured aerosol size distributions for estimating new particle formation rate. A NPF field campaign was carried out from March 7th to Apr. 7th, 2016, in urban Beijing, and a diethylene glycol scanning mobility particle spectrometer equipped with a miniature cylindrical differential mobility analyser was used to measure aerosol size distributions down to ~1 nm. 11 typical NPF events were observed during this period. Measured aerosol size distributions from 1 nm to 10 μm were used to test the new formula and those widely used ones. Previously used formulae that perform well in relatively clean atmosphere where nucleation intensity is not strong were found to underestimate the comparatively high new particle formation rate in urban Beijing because of their underestimation or neglect of the coagulation scavenging effect. Coagulation sink term is the governing component of the estimated formation rate in the observed NPF events in Beijing, and coagulation among newly formed particles contributes a large fraction to the coagulation sink term. Previously reported formation rates in Beijing and in other locations with intense NPF events might be underestimated because the coagulation scavenging effect was not fully considered, e.g., estimated formation rates of 1.5 nm particles in this campaign using the new formula are 1.3 – 4.3 times those estimated using the formula neglecting coagulation among particles in the nucleation mode.

1 Introduction

New particle formation (NPF) is a frequently occurring phenomenon in atmospheric environment. In a typical NPF event, gaseous precursors burst out into particles due to nucleation and lead to a rapid increase in atmospheric aerosol population. Nucleated particles can grow quickly to increase the number concentration of cloud condensation nuclei (Kerminen et al., 2012; Kuang et al., 2009; Leng et al., 2014) and thus has indirect impacts on radiative forcing and global climate (Lohmann & Feichter, 2005). Continuous growth of nucleated particles also provides increasing aerosol surface area for heterogeneous physicochemical processes. NPF studies can trace back to the early 20th century (Aitken, 1911) and NPF events have been observed in various atmospheric environment, e.g., from city to countryside, from desert (Misaki, 1964) to rain forest (Zhou,

2002), from continent to the ocean (Covert et al., 1992), from the equator (Clarke et al., 1998) to polar area (Covert et al., 1996; Park et al., 2004), and from troposphere to stratosphere (Lee et al., 2003).

Formation rate at which the growth flux past a certain diameter is a key parameter to quantitatively describe NPF events. Different formulae have been used to estimate new particle formation rate from measured aerosol size distributions and they mainly originate from two approaches. One is from the definition of nucleation rate (Heisler & Friedlander, 1977; Weber et al., 1996) and the other is a population balance method (Kulmala et al., 2001; Kulmala et al., 2012). Consistency of these two approaches was tested using a numerically simulated NPF event and a relative error of less than 20% was reported (Vuollekoski et al., 2012). The simulated NPF event has a maximum formation rate of less than $1 \text{ cm}^{-3} \text{ s}^{-1}$. However, the reported formation rates in the atmosphere vary in a large scale, e.g., approximately from 10^{-2} to $10^4 \text{ cm}^{-3} \text{ s}^{-1}$ (Kulmala et al., 2004). Suffering from the assumptions made in these two approaches, their validity in describing NPF events with high formation rate needs to be further explored. A high fraction of newly formed particles is scavenged by coagulation before they grow into larger particles. Both approaches potentially underestimate the contribution of coagulation scavenging when calculating formation rate from measurement data. They may perform well in clean atmospheric environment where nucleation intensity is not strong and aerosol concentration is relatively low, i.e., the coagulation scavenging effect is less important.

The effect of coagulation scavenging is more prominent when estimating formation rate of sub-3 nm particles because of their high diffusivities and high concentrations during NPF events. Due to instrument limitations, aerosol size distributions of sub-3 nm particles were not available in many previous NPF field campaigns. Recent developments in diethylene glycol (DEG) condensation particle counters (CPC, Iida et al., 2009; Vanhanen et al., 2011) made it feasible to develop new scanning mobility particle spectrometers (SMPS) for extending aerosol size distribution measurement from ~3 nm down to ~1 nm (Jiang, et al., 2011a; Franchin *et al.*, 2016). These new spectrometers were deployed in atmospheric observations (Jiang, et al., 2011b) and in chamber measurements (Franchin *et al.*, 2016) to study NPF. A miniature cylindrical differential mobility analyser (mini- cyDMA, Cai et al., 2017) was developed to improve the performance of the DEG-SMPS.

In many locations of China, high emissions lead to both high concentrations of gaseous precursors and high atmospheric aerosol concentration. NPF was frequently observed even in megacities such as Beijing and Shanghai (Wu et al., 2007; Kulmala et al., 2016; Wang et al., 2017). In most previous studies, the above population balance method was used to estimate new particle formation rates in China. The reported formation rates of 3 nm particles and larger ones were typically in the range of $1\text{-}10 \text{ cm}^{-3} \text{ s}^{-1}$ (Wang et al., 2013; Leng, et al., 2014; An et al., 2015; Qi et al., 2015). One study in Shanghai reported a rate of 112.4 to 271.0 $\text{cm}^{-3} \text{ s}^{-1}$ for the formation of 1.5 nm particles inferred from a DEG-CPC (Xiao et al., 2015). For these intense NPF events, the above balance approach may underestimate the coagulation scavenging effect and thus lead to underestimation in the reported formation rate. In addition, applying new SMPSs to measure aerosol size distributions down to ~1 nm will help to better quantify the formation rate and its governing factors in typical locations of China.

When estimating new particle formation rates, various particle size ranges were used in previous formulae. The definition approach tries to limit the size range towards the minimum detected diameter (Kuang *et al.*, 2008; Weber, et al., 1996), while studies with the population balance method have used various size ranges. Some studies used the aerosol size distributions from the minimum detected diameter up to 25 nm (Kulmala et al., 2001; Dal Maso et al., 2005; Wu et al., 2007; Wang et al., 2013). Kulmala et al. (2004) recommended the upper size bound as the maximum size that the critical cluster can reach during a short time interval of growth. There are also studies using narrower size range such as from 3 nm to 6 nm (Sihto et al., 2006; Paasonen et al., 2009; Wang et al., 2011; Vuollekoski et al., 2012) and from 1.34 nm to 3 nm (Xiao et al., 2015). In principle, the estimated formation rates may vary when different particle size ranges are used. Assumptions made while deriving these formulae should be fully considered when proposing criteria to choose particle size range.

In this study, a new population balance formula for estimating new particle formation rate was derived from aerosol general dynamic equation to properly account for the effect of coagulation scavenging, especially for analysing intense NPF events. A NPF field campaign was carried out in Beijing. Aerosol size distributions down to ~ 1 nm were measured using the DEG-SMPS equipped with the mini- cyDMA. Data from this campaign and from literature are used to test the new formula and other widely used formulae. Different formulae are compared and their applicability in analysing intense NPF events are addressed. Criteria to choose particle size range for formation rate estimation are proposed and evaluated. Governing components of the new formation rate in Beijing are discussed and compared to those from other locations in the world.

2 Theory

2.1 The new balance formula to estimate formation rate

The new formula based on definition of droplet current and aerosol general dynamic equation (see Appendix A for its derivation) is shown in Eq. (1),

$$J_k = \frac{dN_{[d_k, d_u]}}{dt} + \sum_{d_g=d_k}^{d_u-1} \sum_{d_i=d_{\min}}^{+\infty} \beta_{(i,g)} N_{[d_i, d_{i+1})} N_{[d_g, d_{g+1})} - \frac{1}{2} \sum_{d_g=d_k}^{d_u-1} \sum_{\substack{d_i^3 + d_{j+1}^3 = d_g^3 \\ d_{i+1}^3 + d_j^3 = d_g^3 \\ d_i, d_j \geq d_{\min}}} \beta_{(i,j)} N_{[d_i, d_{i+1})} N_{[d_j, d_{j+1})} + n_u \cdot GR_u \quad (1)$$

where J_k is the formation rate of particles at size d_k ; N is particle number concentration and $N_{[d_k, d_u]}$ is defined as total number concentration of particles ranged from d_k to d_u (particles with diameters of d_u are not accounted for); d_i refers to the lower bound of each measured size bin; $\beta_{(i,g)}$ is the coagulation coefficient when particles with the diameter of d_i collides with particles with the diameter of d_g ; n is particle size distribution function which equals dN/dd_p ; and GR_u is particle growth rate at d_u , i.e., dd_u/dt . d_u is the upper bound of the size range for calculation. d_{\min} is the size of minimum cluster in theory and the lowest size limit of measuring instrument in practice. The last three terms in the right hand side (RHS) of Eq. (1) are coagulation sink term (*CoagSink*), coagulation source term (*CoagSrc*) and condensational growth term, respectively.

89 The two assumptions of Eq. (1) are that (a) transport, dilution, primary emission and other losses except for coagulation loss
 90 in the size range from d_k to d_u are comparatively negligible; (b) when deriving the fourth term in the RHS of Eq. (1), net
 91 coagulation (net result of both formation and scavenging due to coagulation) of any particle larger than d_u with other particles
 92 is negligible. These two assumptions above are also the criteria to determine d_u . The mathematical expression of population
 93 balance in Eq. (1) in the discrete form is illustrated by Fig. 1. Time rate of change of particles at d_k is equal to source minus
 94 sink. Source are the condensational flux into d_k (J_k) and formation due to coagulation among smaller particles/clusters
 95 ($CoagSrc_k$). Sink are the condensational flux out of d_k (J_{k+1}) and scavenging due to coagulation with other particles/clusters
 96 ($CoagSnk_k$). Nucleation rate, I , is defined as J_k when d_k is the size of the critical cluster (nuclei). Equation (1) is obtained by
 97 adding these single population balance equations up from d_k to d_u , converting it from the discrete form into the continuous
 98 form, and approximating J_u with the product of measured n_u and GR_u . Note that Eq. (1) is still an approximate formula of
 99 particle formation rate because $CoagSnk$ and $CoagSrc$ are calculated by size bins and the coagulation effect of particles smaller
 100 than d_{min} is not accounted for. For rigorous mathematical derivation and detailed illustration, please refer to Appendix A.

101 2.2 Previous approaches to estimate formation rate

102 The population balance method proposed in previous study is shown in Eq. (2) (Kulmala et al., 2001; Kulmala et al., 2012),

$$103 \quad J_k = \frac{dN_{[d_k, d_u]}}{dt} + CoagS_m \cdot N_{[d_k, d_u]} + \frac{N_{[d_k, d_u]}}{(d_u - d_k)} \cdot GR_{[d_k, d_u]} \quad (2)$$

104 where coagulation sink, $CoagS_m$, is defined as Eq. (3).

$$105 \quad CoagS_m = \int_0^{+\infty} \beta_{(i,m)} n_i dd_i \quad (3)$$

106 The subscript m corresponds to the representing diameter, d_m , for particles ranged from d_k to d_u . d_m is often estimated as the
 107 geometric mean diameter of d_k and d_u . Equation (1) and (2) look similar because they are both derived from the general dynamic
 108 equation, while their detailed differences are illustrated in Appendix B.

109 The definition approach to calculate new particle formation rate is shown in Eq. (4) (Heisler & Friedlander, 1977; Weber et
 110 al., 1996; Iida et al., 2006; Kuang et al., 2008; Kuang et al., 2012).

$$111 \quad J_k = n_k \cdot GR_k \quad (4)$$

112 Equation (4) focuses on the flux into d_k and is theoretically correct in continuous space of particle diameter. However, when
 113 applying Eq. (4) in practice, size distribution of particles smaller than d_k is required, which is difficult to obtain (See Appendix
 114 B). Usually diameter bins larger than d_k are used to estimate particle formation rate when using the practical expression of Eq.
 115 (4) (e.g., Eq. (9) defined in section 4.3). As illustrated in Fig. 1, such approximation essentially neglects the first three terms
 116 in the RHS of Eq. (1), and may lead to underestimation of particle formation rate because of neglecting the coagulation
 117 scavenging effect especially when analysing intense NPF events.

2.3 Previous formulae for comparison

Equation (5) is a widely used balance formula to estimate formation rate in previous studies (Kulmala et al., 2001; Dal Maso et al., 2005; Wu et al., 2007; Shen et al., 2011; Wang et al., 2013),

$$J_{1.5} = \frac{dN_{[1.5,25]}}{dt} + N_{[1.5,25]} \sum_{d_i=1.5\text{nm}}^{+\infty} \beta_{(i,8)} N_i + \frac{N_{[1.5,25]}}{(25-1.5)\text{nm}} \cdot GR_{[1.5,25]} \quad (5)$$

where N_i is the number concentration of size bin i . Corresponding to those in Eq. (2), d_u is 25 nm and d_m is 8 nm in Eq. (5). By comparing Eq. (5) with Eq. (1), it can be concluded that Eq. (5) estimates $CoagSink$ using a representative $CoagS_m$ and neglects $CoagSrc$. Growth rates in all formulae in section 2.2 were estimated using mode-fitting method suggested in Kulmala et al. (2012).

When calculating $CoagS_m$, particles smaller than d_m (Kulmala et al., 2012) or even d_u are neglected in some previous studies. Corresponding formulae are shown in Eq. (6) and Eq. (7), respectively. The only difference among Eq. (5), Eq (6), and Eq. (7) is the lower bound when calculating $CoagS_m$ in the second term in the RHS of these equations.

$$J_{1.5} = \frac{dN_{[1.5,25]}}{dt} + N_{[1.5,25]} \sum_{d_i=8\text{nm}}^{+\infty} \beta_{(i,8)} N_i + \frac{N_{[1.5,25]}}{(25-1.5)\text{nm}} \cdot GR_{[1.5,25]} \quad (6)$$

$$J_{1.5} = \frac{dN_{[1.5,25]}}{dt} + N_{[1.5,25]} \sum_{d_i=25\text{nm}}^{+\infty} \beta_{(i,8)} N_i + \frac{N_{[1.5,25]}}{(25-1.5)\text{nm}} \cdot GR_{[1.5,25]} \quad (7)$$

The upper bound, d_u , is selected as 6 nm in some recent studies (Sihto et al., 2006; Riipinen et al., 2007; Paasonen et al., 2009; Wang et al., 2011; Vuollekoski et al., 2012; Wang et al., 2015) as shown in Eq. (8).

$$J_{1.5} = \frac{dN_{[1.5,6]}}{dt} + N_{[1.5,6]} \sum_{d_i=1.5\text{nm}}^{+\infty} \beta_{(i,3)} N_i + \frac{N_{[1.5,6]}}{(6-1.5)\text{nm}} \cdot GR_{[1.5,6]} \quad (8)$$

It should be clarified that d_k in Eq. (5)-(8) was usually 3 nm in previous studies due to the absence of sub-3 nm particle size distributions, and d_m in Eq. (8) was 4 nm rather than 3 nm in previous studies because 4 nm is almost the geometrical mean diameter of 3 nm and 6 nm. Particles smaller than 6 nm were neglected when estimating the coagulation sink term in some studies, although its uncertainties will not be discussed here. The expression of condensational growth term, i.e., the third term in the RHS of Eq. (8) varies with studies, however, it does not influence the generality of the following discussion.

In previous studies, several size bins larger than d_k , typically 3 nm, were adopted when using the practical formula of the definition approach (Weber et al., 1996; Kuang et al., 2008), while here the size range from 1.5 nm to 2.5 nm is applied to estimate $J_{1.5}$ as shown in Eq. (9).

$$J_{1.5} = \frac{N_{[1.5,2.5]}}{(2.5-1.5)\text{nm}} \cdot GR_{[1.5,2.5]} \quad (9)$$

3 Experiment

A NPF field campaign was carried out in Beijing. The observation period was from March 7th to April 7th, 2016. The monitoring site locates on the top floor of a four-storey building in the centre of the campus of Tsinghua University. Tsinghua situates in the northwestern urban area of Beijing and the fourth-ring road is ~2 km away to the south of the monitoring site. The site has been a PM_{2.5} monitoring station since 1999 (He et al., 2001; Cao et al., 2014) and there is no tall building nearby. Potential pollution sources around are the three cafeterias on campus that may produce cooking aerosol during meal time, locate ~170 m away on the northeast, ~170 m away on the north, and ~350 m away on the northwest, respectively.

A DEG-SMPS equipped with a mini- cyDMA specially designed for classification of sub-3 nm particles was deployed to measure particles in the size range of 1-5 nm (Cai et al., 2017). A particle size distribution system, including a SMPS with a TSI nano DMA, a SMPS with a TSI long DMA and an aerodynamic particle sizer, was used to measure particles in the size range of 3 nm to 10 μ m in parallel (Liu et al., 2016). Other instruments whose data are not used in this analysis are not listed here.

A C++ program was used to invert particle size distribution from raw counts while incorporating diffusion losses inside the sampling tube, diffusion losses and charging efficiencies of the bipolar neutralizers, penetration efficiencies and transfer functions of DMAs, and detection efficiencies of CPCs (Hagen and Alofs, 1983; Jiang et al., 2011a). Particle density was assumed to be 1.6 g/cm³ according to local observation results (Hu et al., 2012). Accommodation coefficient (i.e., coagulation efficiency) was assumed to be unity and temperature was assumed to be a constant value of 285 K, the average temperature during the observation period.

4 Results and discussion

4.1 Upper size bound for formation rate calculation

New particle formation rates using different upper size bound, d_u of 3 nm, 6 nm, 10 nm and 25 nm were calculated. Since the maximum size that new particles formed by nucleation have reached varies with time, the upper size bound should not be a constant value to minimize the interference of background particles. A varying upper size bound, d_b , was visually determined as the largest size bin in the size range from 3 nm to 25 nm whose frequency density (particle size distribution), $dN/d\log d_p$, was larger than 28,000 #/cm³. Here 28,000 was determined visually according to measured intensity plot of particle size distributions as an approximate boundary for newly formed particles and background particles. The value should be campaign specific or even event specific. Fig. 2(a) indicates that d_b is almost the boundary for particles formed due to nucleation. Estimated $J_{1.5}$ using 20,000 #/cm³ as the boundary differed little from that using 28,000 #/cm³, indicating the estimated $J_{1.5}$ is insensitive to the value for boundary. It is reasonable to regard d_b as a relatively credible value when compared to others. Note

that when using d_b as the upper size bound, dN/dt term of newly formed particles in Eq. (1) is approximated by that of sub-25 nm particles to avoid potential influence of varying size range on particle number concentration.

As shown in Fig 2(b), estimated $J_{1.5}$ using d_b and a constant value of 25 nm as the upper bounds are almost the same (the mean relative error is 2.2%). Maximum difference between these two choices is ~10% which appears before 8:00 when d_b is less than 5 nm and the number concentration of sub-25 nm particles is ~2 times of sub-6 nm particles and ~3 times of sub-3 nm particles. It indicates the influence of non-freshly nucleated particles on estimating $J_{1.5}$ is not important because their comparatively low diffusivities even though their concentration is comparatively high at the beginning of NPF events. Estimated $J_{1.5}$ using d_u of 6 nm and 10 nm are in good consistency with that using d_b before 10:00 (the mean relative errors are 4.8% and 2.6%, respectively). However, when particles formed by nucleation grow beyond, calculated $J_{1.5}$ is underestimated when using 6 nm and 10 nm as the upper bound. For example, the mean relative errors of estimated $J_{1.5}$ using d_u of 6 nm and 10 nm between 10:30 and 15:00 are 18.6% and 12.8%, respectively. When calculating $J_{1.5}$ using 3 nm as d_u , an average 47% underestimation was found for this event.

The reason of underestimation when using smaller d_u can be illustrated by Fig. 2(c). J_u is estimated by $n_u \cdot GR_u$ in Eq. (1). This estimation may be not accurate when d_u is small because the assumption that net coagulation of any particle larger than d_u with other particles is negligible may be violated. As illustrated in the derivation of Eq. (1) in Appendix A, a nearly zero J_u is preferred when using Eq. (1). However, as shown in Fig 2(c), estimated J_3 is still a large fraction compared to $J_{1.5}$, while J_6 and J_{10} are 27.8% and 17.6% of $J_{1.5}$ on average between 10:30 and 15:00, respectively. Although J_u is approximated by $n_u \cdot GR_u$ rather than simply neglected, this approximation may still lead to uncertainties.

Since $J_{1.5}$ estimated by the varying d_b and a constant value of 25 nm are almost the same with an acceptable relative error even under the interference of non-freshly nucleated particles, 25 nm was adopted as the upper bound for calculating J in this study. Since d_u is “proper large” as defined by the two criterions, it is also reasonable to neglect J_u for simplicity. It should be clarified that 25 nm is not necessarily valid for all other studies, because the upper bound should be determined by the two criterions and can be campaign specific. However, it can be concluded that a very small upper bound such as 3 nm is not recommended because particles formed by nucleation surely grow larger than 3 nm in a typical NPF event while intense primary emission of particles around 3 nm is rarely observed in the atmosphere (unless near the emission sources).

4.2 Comparison with previous formulae

Estimated $J_{1.5}$ values using Eq. (1) and Eq. (5)-(9) on March 13th are shown in Fig. 3. d_k , d_u , and d_{min} are 1.5 nm, 25 nm, and 1.3 nm, respectively, when using Eq. (1). It can be concluded that except for Eq. (8), other formulae significantly underestimate $J_{1.5}$ compared to Eq. (1). By comparing contribution of each terms in the RHS of Eq. (1) and Eq. (5)-(9), it was found that the underestimation of formation rates is mainly caused by the underestimation of $CoagSnk$. Equation (9) simply neglects $CoagSnk$

as well as other terms (dN/dt and $CoagSrc$) compared to Eq. (1), so its result is the lowest among six formulae. Equation (5) estimates $CoagSnk$ using an average $CoagS_m$, which lead to underestimation because $CoagS$ at 8 nm happens to be smaller than those at most other diameters in the size range from 1.5 nm to 25 nm, as illustrated in Appendix B. Equation (6) and (7) neglects particles smaller than 8 nm and 25 nm when calculating $CoagS_m$, respectively. Such simplification may be reasonable for relative clean atmosphere where nucleation intensity is not strong, however, these approximations are not suitable for analysing typical NPF events in Beijing where coagulation among nucleation mode particles is a major proportion of $CoagSnk$. $J_{1.5}$ estimated using Eq. (8) agrees well with that estimated using Eq. (1), however, it does not mean that 6 nm serve as a better upper size bound than 25 nm. The agreement between results estimated using Eq. (1) and (8) is because that the estimation of $CoagSnk$ tends to be more accurate when using an average $CoagS_m$ in a narrower size range and the underestimation in this case is coincidentally cancelled out by the overestimation of formation rate caused by neglecting $CoagSrc$.

The importance of coagulation scavenging among newly formed particles due to nucleation is illustrated in Fig. 4. Scavenging due to coagulation with particles smaller than d_p is neglected, as mathematically defined in the formula in Fig. 4(a). $CoagSnk$ increases rapidly with the decrease in d_p rather than maintain an approximately constant value during NPF periods, indicating coagulation among nucleated particles contribute a considerable fraction to $CoagSnk$ in Beijing. The necessity of sub-3 nm particle size distribution is also demonstrated, which means estimated J_3 may also be underestimated due to the absence of sub-3 nm data, as illustrated in Appendix B. Approximation of $CoagSnk$ estimated using a representative $CoagS_m$ is also shown in Fig. 4(b), indicating the underestimation of new particle formation rate when applying Eq. (5) to analyse NPF events in Beijing. However, calculated $CoagSnk$ on a non-NPF event day as well as at non-NPF periods on NPF day is almost unaffected by the coagulation scavenging effect of particles in nucleation mode (smaller than 25 nm), because number concentration of nucleation mode particles at non-NPF time is comparatively low.

4.3 Characteristics of estimated formation rate in Beijing

For NPF events observed in the Beijing campaign, $CoagSnk$ is a governing component of the estimated $J_{1.5}$. Estimated formation rate on March 13th and the four terms in the RHS of Eq. (1), i.e., dN/dt , $CoagSnk$, $CoagSrc$, and the condensational growth term, are shown in Fig. 5. $CoagSnk$ is almost the same with the estimated $J_{1.5}$ in Beijing, while the difference between them is mainly due to dN/dt whose absolute value is comparatively higher at the beginning and the end of the NPF event. The condensational growth term, $n_u \cdot GR_u$, is negligible compared to other terms, which is reasonable since J_u is supposed to be unimportant when determining d_u in Eq. (1). The governing role of $CoagSnk$ in estimated formation rate in Beijing emphasizes the importance of fully considering the coagulation scavenging effect among particles formed by nucleation. Equation (5)-(9) may fit well in relatively clean atmospheric environment where new particle formation rate is comparatively low, such as in Hyytiälä, and agreement of Eq. (8) and Eq. (9) has been reported in a numerically simulated NPF event in which J_3 is less than

1 $\text{cm}^{-3}\text{s}^{-1}$ (Vuollekoski et al., 2012). However, problems appear when applying them in urban Beijing because of underestimating the governing fraction of estimated $J_{1.5}$, i.e., $CoagSnk$.

Coagulation sink, $CoagS$, is not the major reason for the governing role of $CoagSnk$ in Beijing. It is generally considered that the atmosphere in typical urban area in China, such as Beijing, is comparatively polluted. However, observed NPF events mainly occurs on clean days when the air mass comes from north or northwest of Beijing. Mean $\text{PM}_{2.5}$ mass concentration reported by the nearest national monitoring station, Wanliu station, was $10.4 \mu\text{g}/\text{cm}^3$ during all NPF events in this campaign. Aerosol surface area concentration is characterized by Fuchs surface area, A_{Fuchs} (McMurry, 1983), and condensation sink, CS (Kulmala et al., 2001), which are often used to examine the coagulation scavenging effect. The positive correlation between A_{Fuchs} and CS is illustrated in McMurry et al. (2005), while CS can be regarded as the $CoagS$ of sulphuric acid molecules. Fig. 6(a) shows the comparison of A_{Fuchs} and CS in Beijing to those in other locations around the world. A_{Fuchs} and CS during NPF events in this study are higher than those in Hyytiälä, similar to those observed in Boulder, and lower than those in Atlanta, Mexico City, and New Delhi. This indicates coagulation sink in urban Beijing on NPF days is in common range rather than higher than most other places around the world.

Nucleation intensity in urban Beijing, characterized by number concentration of particles larger than 3 nm, is found to be higher than those in Hyytiälä and Atlanta (as shown in Fig. 6(b)). Number concentration of sub-3 nm particles is not accounted for to maintain comparability. Although A_{Fuchs} and $CoagS$ represent the relative importance of the coagulation scavenging effect (McMurry, 1983; Kulmala et al., 2001), it is the $CoagSnk$ that reflects the number of particles lost due to coagulation scavenging in the size range of d_k to d_u . Equation (1) shows that $CoagSnk$ is approximately proportional to the square of particle number concentration. This explains the governing status of $CoagSnk$ in estimated formation rates in urban Beijing in intense NPF events.

Fig. 7 further illustrate the underestimation in new particle formation rates in China due to previously used formulae, especially for Eq. (7) which neglects coagulation among sub-25 nm particles and Eq. (9) which simply neglects net coagulation effect. Mean $J_{1.5}$ estimated in this study using Eq. (1) are 1.2, 2.4, and 6.4 times those estimated using Eq. (5), Eq. (7), and Eq. (9), respectively. Mean J_3 estimated in this study using Eq. (1) are 1.2, 2.0, and 3.3 times those estimated using Eq. (5), Eq. (7), and Eq. (9), respectively. J_3 reported in previous studies in urban Beijing (Wu et al., 2007; Yue et al., 2009; Wang et al., 2013; Wang et al., 2015), Shanghai (Xiao et al., 2015) and Shangdianzi, the regional background station of North China Plain (Shen et al., 2011; Wang et al., 2013), are also shown in Fig. 7. Higher formation rates are anticipated if the coagulation scavenging effect are fully considered when analysing these NPF events. Note that sub-3 nm particles is also accounted when calculating J_3 in this study, while not in previous ones except for the campaign in Shanghai.

4 Conclusions

A new balance formula to estimate new particle formation rate derived from aerosol general dynamic equation was proposed. The new formula estimates the effect of coagulation scavenging better compared to previously used ones. Two criterions in determining the upper bound for calculation were proposed. A NPF campaign in urban Beijing was carried out in spring of 2016. Aerosol size distributions down to ~1 nm was measured and used to test the new formula and those widely used ones in previous studies. It was found that formation rates in urban Beijing are underestimated to different extents in previously used formulae, and the underestimation of the coagulation scavenging effect (corresponding to coagulation sink term) is the major reason. Coagulation among particles in nucleation mode was found to be important when estimating the coagulation scavenging effect in urban Beijing. **Estimated formation rates of 1.5 nm particles in this campaign using the new formula were 1.3 – 4.3 times those estimated using the formula neglecting coagulation among particles in the nucleation mode.** Coagulation sink term is the governing component of the estimated formation rate in urban Beijing. Although higher than those in relative clean atmosphere such as in Hyytiälä, coagulation sink (expressed in the form of Fuchs surface area and condensation sink) in urban Beijing on NPF days is lower than those reported in Atlanta and Mexico City. However, number concentration of particles formed due to nucleation in urban Beijing is comparatively high, which lead to high coagulation loss. Formulae used in previous studies may perform well when describing relative weak NPF events in clean atmosphere, while they underestimate the coagulation scavenging effect when analysing intense NPF events. Formation rates reported in previous studies for urban Beijing and other locations with intense NPF events might be underestimated because of their underestimation or neglect of the coagulation scavenging effect.

Appendix A

Derivation of nucleation rate from aerosol general dynamic equation

Nucleation rate is the rate at which particles grow past the size of the critical cluster (nuclei). However, a more specific and microscopic definition of nucleation rate is needed for any further calculation, and it should be easily and unambiguously transferred into a mathematical expression. Here we adopt the definition based on droplet current (Eq. 10.1, Friedlander, 2000):

$$J_g = \beta_{(1,g-1)} N_1 N_{g-1} - \alpha_g s_g N_g . \quad (A1)$$

Formation rate, J_g , is the excess rate of the passage from $g-1$ (cluster or particle with $g-1$ molecules) to g by condensation over the passage from g to $g-1$ by evaporation. If g is the size of the critical cluster, J_g is defined as nucleation rate, I . N_g is the number concentration of cluster g ; $\beta_{(i,j)}$ is the coagulation coefficient of i and j , and it can be theoretically estimated by diameter of i and j (Eq. 13.56, Seinfeld & Pandis 2006); α_g is the monomer evaporation flux from g ; and s_g is the effective surface area of g for evaporation. Only formation due to condensational growth is considered in the definition of Eq. (A1), while formation

290 due to coagulation of smaller clusters is not taken into account. This is based on the assumption that critical clusters are mainly
 291 formed due to condensational growth of sulfuric acid and other chemical species. The formation of critical cluster by
 292 coagulation does not influence the generality of the following derivation and can be readily incorporated, and it will be clarified
 293 at the end of Appendix A.

294 The other basic equation for the derivation is the general dynamic equation in the discrete form (Eq. 11.3, Friedlander 2000),

$$295 \quad \frac{dN_g}{dt} = \frac{1}{2} \sum_{\substack{i+j=g \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j - \sum_{i=2}^{+\infty} \beta_{(i,g)} N_i N_g + \beta_{(1,g-1)} N_1 N_{g-1} - \beta_{(1,g)} N_1 N_g - \alpha_g s_g N_g + \alpha_{g+1} s_{g+1} N_{g+1}. \quad (A2)$$

296 As shown in Eq. (A2), time rate of change of cluster or particle number concentration, dN_g/dt in the left-hand side (LHS), is
 297 determined by formation due to coagulation of smaller clusters and (or) particles, coagulation scavenging with pre-existing
 298 clusters and particles, condensational growth from $g-1$ and to $g+1$, and evaporation to $g-1$ and from $g+1$, corresponding to the
 299 six terms in the right-hand side (RHS) of Eq. (A2), respectively. The evaporation terms (corresponding to the fifth and sixth
 300 terms in the RHS) may be zero or nearly zero when g is large, however, their exact values have no influence on derivation. An
 301 important assumption to be noted is that meteorological transport, dilution, primary emission of g and other losses (e.g., wall
 302 loss) are not included in Eq. (A2).

303 Notice that the last four terms in the RHS of Eq. (A2) are equal to $J_g - J_{g+1}$ by substituting Eq. (A1) in. Replacing subscript g
 304 with the critical cluster size, k , we have:

$$305 \quad I := J_k = \frac{dN_k}{dt} + \sum_{i=2}^{+\infty} \beta_{(i,k)} N_i N_k - \frac{1}{2} \sum_{\substack{i+j=k \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j + J_{k+1}. \quad (A3)$$

306 The expression of Eq. (A3) is similar to Eq. (A6) in Kuang et al. (2012), which was also obtained using the balance method.
 307 J_{k+1} is usually a relatively large term in Eq. (A3), and it can be accounted for by iteration. Equation (A5) is obtained by
 308 summing Eq. (A3) up from subscript k to $u-1$ as shown in Eq. (A4), where u is the particle size at the upper bound of the
 309 concerned size range.

$$310 \quad \begin{aligned} I - J_{k+1} &= \frac{dN_k}{dt} + \sum_{i=2}^{+\infty} \beta_{(i,k)} N_i N_k - \frac{1}{2} \sum_{\substack{i+j=k \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j \\ J_{k+1} - J_{k+2} &= \frac{dN_{k+1}}{dt} + \sum_{i=2}^{+\infty} \beta_{(i,k+1)} N_i N_{k+1} - \frac{1}{2} \sum_{\substack{i+j=k+1 \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j \\ &\dots = \dots \end{aligned} \quad (A4)$$

$$311 \quad \begin{aligned} J_{u-1} - J_u &= \frac{dN_{u-1}}{dt} + \sum_{i=2}^{+\infty} \beta_{(i,u-1)} N_i N_{u-1} - \frac{1}{2} \sum_{\substack{i+j=u-1 \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j \\ I &= \frac{d \sum_{g=k}^{u-1} N_g}{dt} + \sum_{g=k}^{u-1} \sum_{i=2}^{+\infty} \beta_{(i,g)} N_i N_g - \frac{1}{2} \sum_{g=k}^{u-1} \sum_{\substack{i+j=g \\ i,j \geq 2}} \beta_{(i,j)} N_i N_j + J_u \end{aligned} \quad (A5)$$

In the RHS of Eq. (A5) are the time rate of change of the particle concentration, the coagulation sink term, the coagulation source term and the condensational growth term, respectively. Note that when particle u is large enough, J_u is nearly zero, i.e., $\lim_{u \rightarrow \infty} J_u = 0$, because of their negligible condensational growth and low number concentration compared to those of freshly nucleated small particles. Equation (A6) is obtained by replacing the upper bound, u , with infinite and further simplified by combining the second and third term in the RHS of Eq. (A5).

$$I = -\frac{d \sum_{g=k}^{+\infty} N_g}{dt} + \frac{1}{2} \sum_{g=k}^{+\infty} \sum_{i=k}^{+\infty} \beta_{(i,g)} N_i N_g \quad (\text{A6})$$

Theoretically, Eq. (A6) can be used to estimate I since each term in the RHS can be calculated. However, the validity of Eq. (A6) faces higher risk of violation when applied in real atmosphere due to non-negligible primary emission sources, since Eq. (A6) is a balance equation for the whole aerosol population rather than a limited size range of the nucleation mode. It's both more cautious and efficient to use Eq. (A5) with a proper particle size u and a reasonable estimation of J_u .

When using measured particle size distribution to estimate I , Eq. (A5) has to be converted from the discrete form into the continuous form, i.e., Eq. (A7). Since measured size bins are finite, Eq. (A7) is expressed in the summation form rather than the integration form. Practically, Eq. (A7) is only an estimation of Eq. (A5) because coagulation is calculated by size bins, while particles sizes in each size bin are not exactly the same as the representing diameter, d_g .

$$I = \frac{dN_{[d_k, d_u]}}{dt} + \sum_{d_g=d_k}^{d_{u-1}} \sum_{d_i=d_{\min}}^{+\infty} \beta_{(i,g)} N_{[d_i, d_{i+1}]} N_{[d_g, d_{g+1}]} - \frac{1}{2} \sum_{d_g=d_k}^{d_{u-1}} \sum_{\substack{d_i^3 + d_{j+1}^3 = d_g^3 \\ d_{i+1}^3 + d_j^3 = d_g^3 \\ d_i, d_j \geq d_{\min}}} \beta_{(i,j)} N_{[d_i, d_{i+1}]} N_{[d_j, d_{j+1}]} + J_u \quad (\text{A7})$$

d_{\min} is theoretically the minimum cluster size. The defined relationships, $d_i^3 + d_{j+1}^3 = d_g^3$ and $d_i^3 + d_{j+1}^3 = d_g^3$, are to assure that all particles smaller than d_g are accounted for only once. The third term in the RHS of Eq. A7, i.e., the coagulation source term, is slightly overestimated. Since the coagulation source term is usually much less than the coagulation sink term when d_u is limited (e.g., $d_u < 50$ nm), however, this overestimation can be neglected.

Note that the size bin from d_{u-1} to d_u is denoted by subscript $u-1$, so the upper bound of the size range for calculation is d_u . $N_{[d_k, d_u]}$ is defined as the number concentration in the size range from d_k to d_u (particles with diameters of d_u are not accounted for), corresponding to $\sum_{g=k}^{u-1} N_g$ in the discrete form. d_u is a “proper large” size at which diameter J_u is negligible compared to the sum of the others three terms in the RHS of Eq. (A7). “Proper large” is defined by the following two criterions: the one is d_u shouldn't be too large so that the calculated nucleation rate is non-negligibly affected by transport or primary emissions; the other is d_u shouldn't be too small so that the calculated nucleation rate is underestimated because J_u is still too large to be neglected or to be estimated by growth rate (as illustrated in the following paragraph). These two criterions seem to be contradictory, however, as illustrated in Fig. 2b,

338 calculated nucleation rate is usually not sensitive to the upper bound because J_u decreases rapidly with the increase of d_u since
 339 the freshly nucleated particles are usually in a relatively narrow size range, especially during strong NPF events.
 340 The fourth term in the RHS of Eq. (A7), J_u , is usually so small that it can be simply neglected when d_u is proper large. However,
 341 an approximate term is recommended for better estimation. Here we introduce a sufficient but possibly unnecessary condition
 342 that net coagulation effect between any particle larger than d_u and other particles can be neglected. Define $N_{[d_u, d_u + \Delta d]}|_t$ as
 343 number concentration of particles in a narrow size range from d_u to $d_u + \Delta d$ at time t . After a very short time dt , these particles
 344 grow into the size range from $d_u + dd$ to $d_u + \Delta d + dd$, which is based on the assumption that diameter growth is equal for different
 345 particles in such narrow size and time range, while number concentration remains the same since there is no particle loss.
 346 Particles in the size range from $d_u + \Delta d$ to $+\infty$ at time t grow up to the size range from $d_u + \Delta d + dd$ to $+\infty$, correspondingly. And
 347 since the size range is narrow enough, it's reasonable to assume that concentration of particles is equally distributed in the size
 348 range from d_u to $d_u + \Delta d + dd$, i.e.,

$$349 \frac{N_{[d_i, d_j]}|_{t+dt}}{N_{[d_m, d_n]}|_{t+dt}} = \frac{d_j - d_i}{d_n - d_m}, \text{ for any } d_i, d_j, d_m, d_n \in [d_u, d_u + \Delta d + dd]. \quad (\text{A8})$$

350 Particle size distribution function, n , and growth rate, GR , are defined as Eq. (A9) and A(10), respectively. Equation (A11) is
 351 obtained by combining Eq. (A6), Eq. (A8), Eq. (A9), and Eq. (A10).

$$352 n_u = \frac{dN}{dd} \Big|_{d_u} = \lim_{\Delta d \rightarrow 0} \frac{N_{[d_u, d_u + \Delta d]}}{\Delta d} \quad (\text{A9})$$

$$353 GR_u = \frac{dd}{dt} \Big|_{d_u} \quad (\text{A10})$$

$$\begin{aligned} 354 J_u &= \frac{dN_{[d_u, +\infty]}}{dt} \\ &= \frac{N_{[d_u, d_u + dd]}|_{t+dt} + N_{[d_u + dd, +\infty]}|_{t+dt} - N_{[d_u, +\infty]}|_t}{dt} \\ 355 &= \frac{N_{[d_u, d_u + dd]}|_{t+dt}}{dt} \\ &= \lim_{\Delta d \rightarrow 0} \frac{N_{[d_u, d_u + dd]}|_{t+dt}}{N_{[d_u + dd, d_u + \Delta d + dd]}|_{t+dt} \cdot dt} \cdot N_{[d_u + dd, d_u + \Delta d + dd]}|_{t+dt} \\ 356 &= \lim_{\Delta d \rightarrow 0} \frac{dd}{\Delta d \cdot dt} \cdot N_{[d_u, d_u + \Delta d]}|_t \\ 357 &= n_u \cdot GR_u \quad (\text{A11}) \end{aligned}$$

358 Finally combining Eq. (A7) and Eq. (A11) we can obtain the equation to estimate nucleation rate as Eq. (A12),

$$I = \frac{dN_{[d_k, d_u]}}{dt} + \sum_{d_g=d_k}^{d_u-1} \sum_{d_i=d_{\min}}^{+\infty} \beta_{(i,g)} N_{[d_i, d_{i+1}]} N_{[d_g, d_{g+1}]} - \frac{1}{2} \sum_{d_g=d_k}^{d_u-1} \sum_{\substack{d_i^3 + d_{j+1}^3 = d_g^3 \\ d_{i+1}^3 + d_j^3 = d_g^3 \\ d_i, d_j \geq d_{\min}}} \beta_{(i,j)} N_{[d_i, d_{i+1}]} N_{[d_j, d_{j+1}]} + n_u \cdot GR_u. \quad (\text{A12})$$

The first term in the RHS of Eq. (A12) is the change in the number concentration of particles ranged from d_k to d_u . The second and third terms are particle loss to coagulation scavenging and particle formation by coagulation, named as coagulation sink term (*CoagSnk*) and coagulation source term (*CoagSrc*), respectively (Kuang et al, 2012). The fourth term is the condensational growth term, which is an approximation of the formation rate, J_u . This balance formula derived from aerosol general dynamic equation can also be expressed as Eq. (A13).

$$I = \frac{dN_{[d_k, d_u]}}{dt} + CoagSnk - CoagSrc + n_u \cdot GR_u \quad (\text{A13})$$

When applying Eq. (A12) in practice, d_k is usually the assumed size of the critical nuclei (or the lowest size limit of instrument, corresponding to formation rate, J_k , rather than nucleation rate, I). The dN/dt term can be obtained either by differentiating between adjacent time bins or fitting in a continuous time period. *CoagSnk* and *CoagSrc* can be directly calculated from particle size distribution, where d_{\min} is the minimum detected particle diameter. If formation by coagulation of smaller clusters is also included in the definition of nucleation rate, calculation of *CoagSrc* (the third term in the RHS of equation A(12)) should begin with d_{k+1} instead of d_k , which usually affects little since the difference is only a size bin and the whole *CoagSrc* is usually a minor term of J in atmosphere environment. Growth rate can be estimated by different methods (Weber et al., 1996; Weber et al., 1997; Kulmala et al., 2012; Lehtipalo et al., 2014), or the growth term can be simply neglect if d_u is proper large. It should be clarified that the formation rate calculated using Eq. (A12) may be underestimated because coagulation scavenging by particles and clusters smaller than d_{\min} is neglected due to the limitation of measuring instruments. As illustrated in Fig. 6(a), *CoagSnk* calculated using d_p larger than 3 nm is ~ 89.1% of that using d_p larger than 1.5 nm. It could be inferred that the calculated J_3 was slightly underestimated in some previous studies lacking size distribution for sub-3 nm particles. While in this study, measured particles down to 1.3 nm are accounted for when calculating $J_{1.5}$ and J_3 . Neglecting coagulation between clusters may also have a non-negligible effect on the calculated results (McMurry 1983), which calls for measurement of major molecular clusters participating in nucleation if more accurate formation rate is to be obtained.

Appendix B

Relationships with previous approaches

Since the new balance approach proposed in this study is based on aerosol general dynamic equation with a reasonable assumption that net coagulation of any particle larger than the “proper large” upper bound, d_u , and other particles can be neglected, its inner relationships with former approaches can be elucidated by making additional assumptions and approximations.

Formation rate is defined as the flux that particles grow pass through the given size, and can be expressed as Eq. (B1), where k is the number of molecules contained by the particle (Heisler & Friedlander, 1977; Weber et al., 1996; Kuang et al., 2008; Kuang et al., 2012). Note that Eq. (B1) is valid only when it is in the continuous space of particle diameter, while a more accurate expression in the discrete form is shown as Eq. (B2).

$$J_k = n_k \cdot GR_k \quad (B1)$$

$$J_k = n_{k-1} \cdot GR_{k-1} \quad (B2)$$

Eq. (B2) is believed to be theoretically correct since the only condensational flux into d_k is the growth of smaller clusters or particles with diameter of d_{k-1} . Although in similar expression with Eq. (A11), Eq. (B2) focuses on the flux into rather than out of the size bin for calculation, and there's no need to account for coagulation scavenging, as illustrated in Fig. 1.

A theoretical expression of GR proposed in previous study is shown as Eq. (B3), where α is herein the coagulation efficiency (fraction of collisions that successfully result in coagulation), V_1 is the volume increment when adding a single gaseous precursor, and v is the mean thermal velocity of the gaseous precursor (Weber et al., 1996). Here we update the equation by considering different chemical species and describing coagulation by β , as shown in Eq. (B4). The subscript c denotes different chemical species of monomers participating in the condensational growth of cluster $k-1$, and N_{1c} is their corresponding number concentration. Coagulation efficiency is included in each $\beta_{(1c,k)}$ (Eq. 13.56, Seinfeld & Pandis 2006).

$$GR_k = \frac{\alpha V_1 N_1 v}{2} \quad (B3)$$

$$GR_{k-1} = \frac{\sum \beta_{(1c,k-1)} N_{1c} N_{k-1}}{n_{k-1}} \quad (B4)$$

Eq. (B2) is theoretically correct, however, it faces difficulties when applying in practice, since n_{k-1} is obtained by approximation over some size range around d_k rather than the true frequency density at cluster $k-1$, dN_{k-1}/dd_{k-1} . Moreover, because size distribution smaller than d_k is difficult to obtain, the size range for estimation is usually larger than d_k . For example, the formula to estimate J_3 using nano-SMPS data in Kuang et al. (2008) is shown as Eq. (B5). Although Eq. (B5) seems to be an estimation of Eq. (B2), they are essentially two different equations. This is because the measured particle number concentration in the size range for calculation, i.e., N_{3-6} in Eq. (B5), has been affected by coagulation. By comparing with Eq. (A13), it can be concluded that dN/dt , $CoagSink$ and $CoagSrc$ are simply neglected in Eq. (B5), while Eq.(B2) does not suffer from this problem by its definition.

$$J_3 \approx \frac{N_{3-6}}{3 \text{ nm}} \cdot GR_{1-3} \quad (B5)$$

There are also problems in estimating GR_{k-1} . Equation (B4) is only a theoretical formula, since it is nearly impossible to determine all the chemical species contributing to nucleation and their corresponding coagulation coefficients in the complicated atmospheric environment. GR calculated by sulfuric acid itself using Eq. (B3) may lead to underestimation (Kuang et al., 2010), while uncertainties also exist in the approaches which fit particles size distribution to obtain GR (Kulmala et al.,

2012; Lehtipalo et al., 2014) because the effect of coagulation on measured size distribution is also neglected. So conclusively, Eq. (B2) is considered to be theoretically correct, however, it's not recommend to be applied for analyzing NPF events with high coagulation scavenging.

The other approach is a balance method based on a macroscopic point of view shown as Eq. (B6) (Kulmala et al., 2001; Kulmala et al., 2004), and here we adopt the equation in the most recent paper (Kulmala et al, 2012). Usually d_m is the geometric mean diameter of d_k and d_u . However, coagulation between any particle smaller than d_m or even d_u with another particle (with any size) is sometimes neglected when it comes to calculation, such as the formula suggested in Kulmala et al (2012) shown as Eq. (B7).

$$J_k = \frac{dN_{[d_k, d_u]}}{dt} + CoagS_m \cdot N_{[d_k, d_u]} + \frac{N_{[d_k, d_u]}}{(d_u - d_k)} \cdot GR_{[d_k, d_u]} \quad (B6)$$

$$CoagS'_m = \sum_{d_i=d_m}^{+\infty} \beta_{(i,m)} N_i \quad (B7)$$

Eq. (B6) appears similar to Eq. (A13) since they both originate from the population balance method, however, there are some differences between them.

Firstly, the upper bound of particle size in Eq. (B6), d_u , is lack of strict definition and discussion. As discussed in Appendix A, d_u should be decided by the two criterions that effects of transport and primary emission are negligible and the condensational growth term, J_u , is relative small compared to J_k . The upper bound of 25 nm is usually reasonable since high concentration of particle formed by nucleation predominates the coagulation sink term during strong new particle formation time, while the upper bound of 6 nm may lead to underestimation when freshly formed particles grow beyond, as discussed in the main text. Secondly, scavenging by coagulation with particles smaller than d_m is not included if using Eq. (B7) to calculate $CoagS$. As shown in Fig. B1, $CoagS$ is always larger than $CoagS'$, and their difference increases as d_m increases. $CoagS'_{8nm}$ is ~31% of $CoagS_{8nm}$, indicating a large amount of underestimation when using Eq. (B7). Note that Eq. (3) and the approximation formula (estimated with condensation sink) proposed by Lehtinen et al. (2007) does not suffer from this problem.

Thirdly, the second term in the RHS of Eq. (B6) is not always a reasonable approximation of $CoagSnk$ in Eq. (A12) and Eq. (A13). Theoretically, the relationship between $CoagSnk$ and $CoagS$ is shown as Eq. (B8), while $CoagS_m$ is chosen as the representative value when estimating J using Eq. (B6).

$$CoagSnk = \sum_{d_g=d_k}^{d_u} CoagS_g \cdot N_g \quad (B8)$$

However, neither is $CoagS$ a relatively constant value versus particle diameter nor is $CoagS_m$ the mean value of $CoagS$ in calculated size range from d_k to d_u . As illustrated in Fig. B1, coagulation coefficient with 8 nm particles decreases rapidly with the increase in d_i when particle is smaller than 8 nm. The minimum value of $\beta_{(d_i, 8nm)}$ appears at d_i around 8 nm because particles with similar thermal velocities are more difficult to collide with each other. The calculated $CoagS'$ during a strong

446 NPF event on Mar. 27th, 2016 appears monotonously decreasing with the increase of d_m , while the calculated $CoagS$ has a
 447 minimum value at 6.7 nm because $CoagS$ is mainly attributed to nucleation mode particles during NPF events. In this example,
 448 $CoagS_{8nm}$ and $CoagS'_{8nm}$ are ~22.6% and ~7.2% of $CoagS_{1.5nm}$, respectively, indicating non-negligible underestimation of
 449 coagulation sink term as well as nucleation rate when using a constant $CoagS_m$ instead of a varying value (as a function of
 450 particle diameter).
 451 Fourthly, particle formation by coagulation is neglected in Eq. (B6). The absence of $CoagSrc$ will lead to an overestimation of
 452 nucleation rate. However, it sometimes coincidentally cancels out with the underestimation caused by using $CoagS_m$ to
 453 approximate $CoagSrc$, as discussed in the main text.
 454 Fifthly, the growth term in Eq. (B6) is estimated over the whole size range from d_k to d_u , while in Eq. (A12) it is mathematically
 455 restricted at the upper bound, d_u . n_u is usually smaller than mean value in the size range from d_k to d_u during a NPF event, and
 456 recent work have revealed that the observed GR is size dependent (Kuang et al., 2012; Kulmala et al., 2013; Xiao et al., 2015).
 457 For example, as shown in Fig. B2, GR varies with time in the NPF event on Apr. 3rd, 2016, and was linearly fitted in different
 458 diameter ranges. The mean GR of particles ranged from 2 nm to 25 nm is ~7.47 nm/h, while GR_{25} is ~10.86 nm/h. At 11:30
 459 on Apr. 3rd, n_{25} ($dN/d\log d_p$ at 25 nm) is 164 #/cm³, while the mean n of particles ranged from 2 nm to 25 nm is 4755 #/cm³.
 460 The calculated condensational growth term in Eq. (B6) is ~20 times of that in Eq. (A12).
 461 In relatively clean environment with weak NPF events, Eq. (B6) may work well since the calculated J_k is mainly predominated
 462 by dN/dt . However, when number concentration of aerosol formed by nucleation and (or) background aerosol is high, i.e.,
 463 when $CoagSnk$ is the major component of J_k , Eq. (B6) underestimates the formation rate (and nucleation rate) due to
 464 underestimation of the coagulation scavenging effect.

465

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469 References

- 470 Aitken, J.: On some nuclei of cloudy nucleation, Proceedings of the Royal Society of Edinburgh, XXXI, 495-511, 1911.
 471 An, J., Wang, H., Shen, L., Zhu, B., Zou, J., Gao, J., and Kang, H.: Characteristics of new particle formation events in Nanjing, China: Effect
 472 of water-soluble ions, Atmospheric Environment, 108, 32-40, doi:10.1016/j.atmosenv.2015.01.038, 2015.
 473 Cai, R., Chen, D.-R., Hao, J., and Jiang, J.: A miniature cylindrical differential mobility analyzer for sub-3 nm particle sizing, Journal of
 474 Aerosol Science, 106, 111-119, doi:10.1016/j.jaerosci.2017.01.004, 2017.
 475 Cao, C., Jiang, W., Wang, B., Fang, J., Lang, J., Tian, G., Jiang, J., and Zhu, T. F.: Inhalable microorganisms in Beijing's PM2.5 and PM10
 476 pollutants during a severe smog event, Environmental science & technology, 48, 1499-1507, doi:10.1021/es4048472, 2014.

Clarke, A. D., Davis, D., Kapustin, V. N., Eisele, F., Chen, G., Paluch, I., Lenschow, D., Bandy, A. R., Thornton, D., Moore, K., Mauldin, L., Tanner, D., Litchy, M., Carroll, M. A., Collins, J., and Albercook, G.: Particle nucleation in the tropical boundary layer and its coupling to marine sulfur sources, *Science*, 282, 89-92, doi:10.1126/science.282.5386.89, 1998.

Covert, D. S., Kapustin, V. N., Quinn, P. K., and Bates, T. S.: New particle formation in the marine boundary layer, *Journal of Geophysical Research*, 97, 20518-20589, 1992.

Covert, D. S., Wiedensohler, A., Aalto, P., Heintzenberg, J., McMurry, P. H., and Leck, C.: Aerosol number size distributions from 3 to 500 nm diameter in the arctic marine boundary layer during summer and autumn, *Tellus B*, 48, doi:10.1034/j.1600-0889.1996.t01-1-00005.x, 1996.

Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P., and Lehtinen, K. E.: Formation and growth of fresh atmospheric aerosols: eight years of aerosol size distribution data from SMEAR II, Hyytiälä, Finland, *Boreal Environment Research*, 10, 323-336, 2005.

Franchin, A., Downard, A., Kangasluoma, J., Nieminen, T., Lehtipalo, K., Steiner, G., Manninen, E. H., Petäjä, T., Flagan, R. C., and Kulmala, M.: A new high-transmission inlet for the Caltech nano-RDMA for size distribution measurements of sub-3 nm ions at ambient concentrations, *Atmospheric Measurement Techniques*, 9, 2709-2720, doi:10.5194/amt-9-2709-2016, 2016.

Friedlander, S. K.: Smoke, dust and haze, 2nd ed., Topics in Chemical Engineering, edited by: Gubbins, K. E., Oxford University Press, New York, 2000.

Hagen, D. E., and Alofs, D. J.: Linear Inversion Method to Obtain aerosol Size Distributions from Measurements with a Differential Mobility Analyzer, *Aerosol Science and Technology*, 2, 465-475, 1983.

He, K., Yang, F., Ma, Y., Zhang, Q., Yao, X., Chan, C. K., Cadel, S., Chan, T., and Mulawa, P.: The characteristics of PM_{2.5} in Beijing, China, *Atmospheric Environment*, 35, 4959-4970, 2001.

Heisler, S. L., and Friedlander, S. K.: Gas-to-particle conversion in photochemical smog: Aerosol growth laws and mechanisms for organics, *Atmospheric Environment*, 11, 157-168, doi:10.1016/0004-6981(77)90220-7, 1977.

Hu, M., Peng, J., Sun, K., Yue, D., Guo, S., Wiedensohler, A., and Wu, Z.: Estimation of size-resolved ambient particle density based on the measurement of aerosol number, mass, and chemical size distributions in the winter in Beijing, *Environmental science & technology*, 46, 9941-9947, doi:10.1021/es204073t, 2012.

Iida, K., Stolzenburg, M., McMurry, P., Dunn, M. J., Smith, J. N., Eisele, F., and Keady, P.: Contribution of ion-induced nucleation to new particle formation: Methodology and its application to atmospheric observations in Boulder, Colorado, *Journal of Geophysical Research: Atmospheres*, 111, D23201, doi:10.1029/2006jd007167, 2006.

Iida, K., Stolzenburg, M. R., McMurry, P. H., and Smith, J. N.: Estimating nanoparticle growth rates from size-dependent charged fractions: Analysis of new particle formation events in Mexico City, *Journal of Geophysical Research: Atmospheres*, 113, doi:10.1029/2007jd009260, 2008.

Iida, K., Stolzenburg, M. R., and McMurry, P. H.: Effect of Working Fluid on Sub-2 nm Particle Detection with a Laminar Flow Ultrafine Condensation Particle Counter, *Aerosol Science and Technology*, 43, 81-96, doi:10.1080/02786820802488194, 2009.

Jiang, J., Chen, M., Kuang, C., Attoui, M., and McMurry, P. H.: Electrical Mobility Spectrometer Using a Diethylene Glycol Condensation Particle Counter for Measurement of Aerosol Size Distributions Down to 1 nm, *Aerosol Science and Technology*, 45, 510-521, doi:10.1080/02786826.2010.547538, 2011a.

Jiang, J., Zhao, J., Chen, M., Eisele, F. L., Scheckman, J., Williams, B. J., Kuang, C., and McMurry, P. H.: First Measurements of Neutral Atmospheric Cluster and 1–2 nm Particle Number Size Distributions During Nucleation Events, *Aerosol Science and Technology*, 45, ii-v, doi:10.1080/02786826.2010.546817, 2011b.

Kerminen, V.-M., Paramonov, M., Anttila, T., Riipinen, I., Fountoukis, C., Korhonen, H., Asmi, E., Laakso, L., Lihavainen, H., Swietlicki, E., Svenningsson, B., Asmi, A., Pandis, S. N., Kulmala, M., and Petäjä, T.: Cloud condensation nuclei production associated with atmospheric nucleation: a synthesis based on existing literature and new results, *Atmospheric Chemistry and Physics*, 12, 12037-12059, doi:10.5194/acp-12-12037-2012, 2012.

Kuang, C., Chen, M., Zhao, J., Smith, J., McMurry, P. H., and Wang, J.: Size and time-resolved growth rate measurements of 1 to 5 nm freshly formed atmospheric nuclei, *Atmospheric Chemistry and Physics*, 12, 3573-3589, doi:10.5194/acp-12-3573-2012, 2012.

Kuang, C., McMurry, P. H., and McCormick, A. V.: Determination of cloud condensation nuclei production from measured new particle formation events, *Geophysical Research Letters*, 36, doi:10.1029/2009gl037584, 2009.

524 Kuang, C., McMurry, P. H., McCormick, A. V., and Eisele, F. L.: Dependence of nucleation rates on sulfuric acid vapor concentration in
525 diverse atmospheric locations, *Journal of Geophysical Research*, 113, doi:10.1029/2007jd0009253, 2008.

526 Kuang, C., Riipinen, I., Sihto, S. L., Kulmala, M., McCormick, A. V., and McMurry, P. H.: An improved criterion for new particle formation
527 in diverse atmospheric environments, *Atmospheric Chemistry and Physics*, 10, 8469-8480, doi:10.5194/acp-10-8469-2010, 2010.

528 Kulmala, M., Dal Maso, M., Mäkelä, J. M., Pirjola, L., Väkevä, M., Aalto, P., Miikkulainen, P., Hämeri, K., and O'dowd, C. D.: On the
529 formation, growth and composition of nucleation mode particles, *Tellus*, 53, 479-490, 2001.

530 Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä, M., Schobesberger, S., Rantala,
531 P., Franchin, A., Jokinen, T., Jarvinen, E., Äijälä, M., Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J.,
532 Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., 3rd, Duplissy, J., Vehkamäki, H., Bäck, J.,
533 Kortelainen, A., Riipinen, I., Kurtén, T., Johnston, M. V., Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E., Laaksonen, A.,
534 Kerminen, V.-M., and Worsnop, D. R.: Direct observations of atmospheric aerosol nucleation, *Science*, 339, 943-946,
535 doi:10.1126/science.1227385, 2013.

536 Kulmala, M., Petäjä, T., Kerminen, V. M., Kujansuu, J., Ruuskanen, T., Ding, A. J., Nie, W., Hu, M., Wang, Z., Wu, Z., Wang, L., and
537 Worsnop, D. R.: On secondary new particle formation in China, *Frontiers of Environmental Science & Engineering*, 10, 191-200,
538 doi:10.1007/s11783-016-0850-1, 2016.

539 Kulmala, M., Petäjä, T., Mönkkönen, P., Koponen, I. K., Dal Maso, M., Aalto, P. P., Lehtinen, K. E., and Kerminen, V. M.: On the growth
540 of nucleation mode particles: source rates of condensable vapor in polluted and clean environments, *Atmospheric Chemistry and
541 Physics*, 5, 409-416, 2005.

542 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P. P., Junninen, H., Paasonen, P.,
543 Riipinen, I., Lehtinen, K. E., Laaksonen, A., and Kerminen, V. M.: Measurement of the nucleation of atmospheric aerosol particles,
544 *Nature protocols*, 7, 1651-1667, doi:10.1038/nprot.2012.091, 2012.

545 Kulmala, M., Vehkamäki, H., Petäjä, T., Dal Maso, M., Lauri, A., Kerminen, V. M., Birmili, W., and McMurry, P. H.: Formation and growth
546 rates of ultrafine atmospheric particles: a review of observations, *Journal of Aerosol Science*, 35, 143-176,
547 doi:10.1016/j.jaerosci.2003.10.003, 2004.

548 Lee, S.-H., Reeves, J. M., Wilson, J. C., Hunton, D. E., Viggiano, A. A., Miller, T. M., Ballenthin, J. O., and Lait, L. R.: Particle formation
549 by ion nucleation in the upper troposphere and lower stratosphere, *Science*, 301, 1886-1889, doi:10.1126/science.1087236, 2003.

550 Lehtinen, K. E. J., Dal Maso, M., Kulmala, M., and Kerminen, V.-M.: Estimating nucleation rates from apparent particle formation rates
551 and vice versa: Revised formulation of the Kerminen–Kulmala equation, *Journal of Aerosol Science*, 38, 988-994,
552 doi:10.1016/j.jaerosci.2007.06.009, 2007.

553 Lehtipalo, K., Leppä, J., Kontkanen, J., Kangasluoma, J., Franchin, A., Wimmer, D., Schobesberger, S., Junninen, H., Petäjä, T., Sipilä, M.,
554 Mikkilä, J., Vanhanen, J., Worsnop, D. R., and Kulmala, M.: Methods for determining particle size distribution and growth rates
555 between 1 and 3 nm using the Particle Size Magnifier, *Boreal Environment Research*, 19, 215-236, 2014.

556 Leng, C., Zhang, Q., Tao, J., Zhang, H., Zhang, D., Xu, C., Li, X., Kong, L., Cheng, T., Zhang, R., Yang, X., Chen, J., Qiao, L., Lou, S.,
557 Wang, H., and Chen, C.: Impacts of new particle formation on aerosol cloud condensation nuclei (CCN) activity in Shanghai: case
558 study, *Atmospheric Chemistry and Physics*, 14, 11353-11365, doi:10.5194/acp-14-11353-2014, 2014.

559 Liu, J., Jiang, J., Zhang, Q., Deng, J., and Hao, J.: A spectrometer for measuring particle size distributions in the range of 3 nm to 10 µm,
560 *Frontiers of Environmental Science & Engineering*, 10, 63-72, doi:10.1007/s11783-014-0754-x, 2016.

561 Lohmann, U., and Feichter, J.: Global indirect aerosol effects: a review, *Atmospheric Chemistry and Physics*, 5, 715–737, doi:10.5194/acp-
562 5-715-2005, 2005.

563 Manninen, E. H., Nieminen, T., Riipinen, I., Yli-Juuti, T., Gagné, S., Asmi, E., Aalto, P. P., Petäjä, T., Kerminen, V.-M., and Kulmala, M.:
564 Charged and total particle formation and growth rates during EUCAARI 2007 campaign in Hyytiälä, *Atmospheric Chemistry and
565 Physics*, 9, 4077-4089, doi:10.5194/acp-9-4077-2009, 2009.

566 McMurry, P. H.: New particle formation in the presence of an aerosol: rates, time scales, and sub-0.01 µm size distributions *Journal of
567 Colloid and Interface Science*, 95, 72-80, 1983.

568 McMurry, P. H., Fink, M., Sakurai, H., Stolzenburg, M. R., Mauldin, R. L., Smith, J., Eisele, F., Moore, K., Sjostedt, S., Tanner, D., Huey,
569 L. G., Nowak, J. B., Edgerton, E., and Voisin, D.: A criterion for new particle formation in the sulfur-rich Atlanta atmosphere,
570 *Journal of Geophysical Research*, 110, doi:10.1029/2005jd005901, 2005.

571 Misaki, M.: Mobility spectrums of large Ions in the new Mexico semidesert, *Journal of Geophysical Research*, 69, 3309-3318, 1964.

572 Paasonen, P., Sihto, S.-L., Nieminen, T., Vuollekoski, H., Riipinen, I., Plaß-Dülmer, C., Berresheim, H., Birmili, W., and Kulmala, M.:
573 Connection between new particle formation and sulphuric acid at Hohenpeissenberg (Germany) including the influence of organic
574 compounds, *Boreal Environment Research*, 14, 616-629, 2009.

575 Park, J., Sakurai, H., Vollmers, K., and McMurry, P. H.: Aerosol size distributions measured at the South Pole during ISCAT, *Atmospheric*
576 *Environment*, 38, 5493-5500, doi:10.1016/j.atmosenv.2002.12.001, 2004.

577 Qi, X. M., Ding, A. J., Nie, W., Petäjä, T., Kerminen, V.-M., Herrmann, E., Xie, Y. N., Zheng, L. F., Manninen, H., Aalto, P., Sun, J. N.,
578 Xu, Z. N., Chi, X. G., Huang, X., Boy, M., Virkkula, A., Yang, X.-Q., Fu, C. B., and Kulmala, M.: Aerosol size distribution and
579 new particle formation in the western Yangtze River Delta of China: 2 years of measurements at the SORPES station, *Atmospheric*
580 *Chemistry and Physics*, 15, 12445-12464, doi:10.5194/acp-15-12445-2015, 2015.

581 Riipinen, I., Sihto, S.-L., Kulmala, M., Arnold, F., Dal Maso, M., Birmili, W., Saarnio, K., Teinilä, K., Kerminen, V.-M., Laaksonen, A.,
582 and Lehtinen, K. E. J.: Connections between atmospheric sulphuric acid and new particle formation during QUEST III-IV
583 campaigns in Heidelberg and Hyytiälä, *Atmospheric Chemistry and Physics*, 7, 1899-1914, doi:10.5194/acpd-6-10837-2006,
584 2007.

585 Seinfeld, J. H., and Pandis, S. N.: *Atmospheric Chemistry and Physics*, 2nd ed., John Wiley & Sons, Inc., New Jersey, 2006.

586 Shen, X. J., Sun, J. Y., Zhang, Y. M., Wehner, B., Nowak, A., Tuch, T., Zhang, X. C., Wang, T. T., Zhou, H. G., Zhang, X. L., Dong, F.,
587 Birmili, W., and Wiedensohler, A.: First long-term study of particle number size distributions and new particle formation events
588 of regional aerosol in the North China Plain, *Atmospheric Chemistry and Physics*, 11, 1565-1580, doi:10.5194/acp-11-1565-2011,
589 2011.

590 Sihto, S.-L., Kulmala, M., Kerminen, V.-M., Dal Maso, M., Petäjä, T., Riipinen, I., Korhonen, H., Arnold, F., Janson, R., Boy, M.,
591 Laaksonen, A., and Lehtinen, K. E.: Atmospheric sulphuric acid and aerosol formation: implications from atmospheric
592 measurements for nucleation and early growth mechanisms, *Atmospheric Chemistry and Physics*, 6, 4079-4091, doi:10.5194/acp-
593 6-4079-2006, 2006.

594 Vanhanen, J., Mikkilä, J., Lehtipalo, K., Sipilä, M., Manninen, H. E., Siivola, E., Petäjä, T., and Kulmala, M.: Particle Size Magnifier for
595 Nano-CN Detection, *Aerosol Science and Technology*, 45, 533-542, doi:10.1080/02786826.2010.547889, 2011.

596 Vuollekoski, H., Sihto, S.-L., Kerminen, V.-M., Kulmala, M., and Lehtinen, K. E. J.: A numerical comparison of different methods for
597 determining the particle formation rate, *Atmospheric Chemistry and Physics*, 12, 2289-2295, doi:10.5194/acp-12-2289-2012,
598 2012.

599 Wang, Z. B., Hu, M., Pei, X. Y., Zhang, R. Y., Paasonen, P., Zheng, J., Yue, D. L., Wu, Z. J., Boy, M., and Wiedensohler, A.: Connection
600 of organics to atmospheric new particle formation and growth at an urban site of Beijing, *Atmospheric Environment*, 103, 7-17,
601 doi:10.1016/j.atmosenv.2014.11.069, 2015.

602 Wang, Z. B., Hu, M., Sun, J. Y., Wu, Z. J., Yue, D. L., Shen, X. J., Zhang, Y. M., Pei, X. Y., Cheng, Y. F., and Wiedensohler, A.:
603 Characteristics of regional new particle formation in urban and regional background environments in the North China Plain,
604 *Atmospheric Chemistry and Physics*, 13, 12495-12506, doi:10.5194/acp-13-12495-2013, 2013.

605 Wang, Z. B., Hu, M., Yue, D. L., Zheng, J., Zhang, R. Y., Wiedensohler, A., Wu, Z. J., Nieminen, T., and Boy, M.: Evaluation on the role
606 of sulfuric acid in the mechanisms of new particle formation for Beijing case, *Atmospheric Chemistry and Physics*, 11, 12663-
607 12671, doi:10.5194/acp-11-12663-2011, 2011.

608 Wang, Z., Wu, Z., Yue, D., Shang, D., Guo, S., Sun, J., Ding, A., Wang, L., Jiang, J., Guo, H., Gao, J., Cheung, H. C., Morawska, L.,
609 Keywood, M., and Hu, M.: New particle formation in China: Current knowledge and further directions, *The Science of the total*
610 *environment*, 577, 258-266, doi:10.1016/j.scitotenv.2016.10.177, 2017.

611 Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measured atmospheric new particleformation rates:
612 implications for nucleation mechanisms, *Chemical Engineering Communications*, 151, 53-64, doi:10.1080/00986449608936541,
613 1996.

614 Weber, R. J., Marti, J. J., McMurry, P. H., Eisele, F. L., Tanner, D. J., and Jefferson, A.: Measurements of new particle formation and
615 ultrafine particle growth rates at a clean continental site, *Journal of Geophysical Research: Atmospheres*, 102, 4375-4385,
616 doi:10.1029/96jd03656, 1997.

617 Wu, Z., Hu, M., Liu, S., Wehner, B., Bauer, S., Maßling, A., Wiedensohler, A., Petäjä, T., Dal Maso, M., and Kulmala, M.: New particle
618 formation in Beijing, China: Statistical analysis of a 1-year data set, *Journal of Geophysical Research*, 112,
619 doi:10.1029/2006jd007406, 2007.

620 Xiao, S., Wang, M. Y., Yao, L., Kulmala, M., Zhou, B., Yang, X., Chen, J. M., Wang, D. F., Fu, Q. Y., Worsnop, D. R., and Wang, L.:
621 Strong atmospheric new particle formation in winter in urban Shanghai, China, *Atmospheric Chemistry and Physics*, 15, 1769-
622 1781, doi:10.5194/acp-15-1769-2015, 2015.

623 Yue, D., Hu, M., Wu, Z., Wang, Z., Guo, S., Wehner, B., Nowak, A., Achtert, P., Wiedensohler, A., Jung, J., Kim, Y. J., and Liu, S.:
624 Characteristics of aerosol size distributions and new particle formation in the summer in Beijing, *Journal of Geophysical Research*,
625 114, doi:10.1029/2008jd010894, 2009.

626 Zhou, J.: Submicrometer aerosol particle size distribution and hygroscopic growth measured in the Amazon rain forest during the wet season,
627 *Journal of Geophysical Research*, 107, doi:10.1029/2000jd000203, 2002.

628

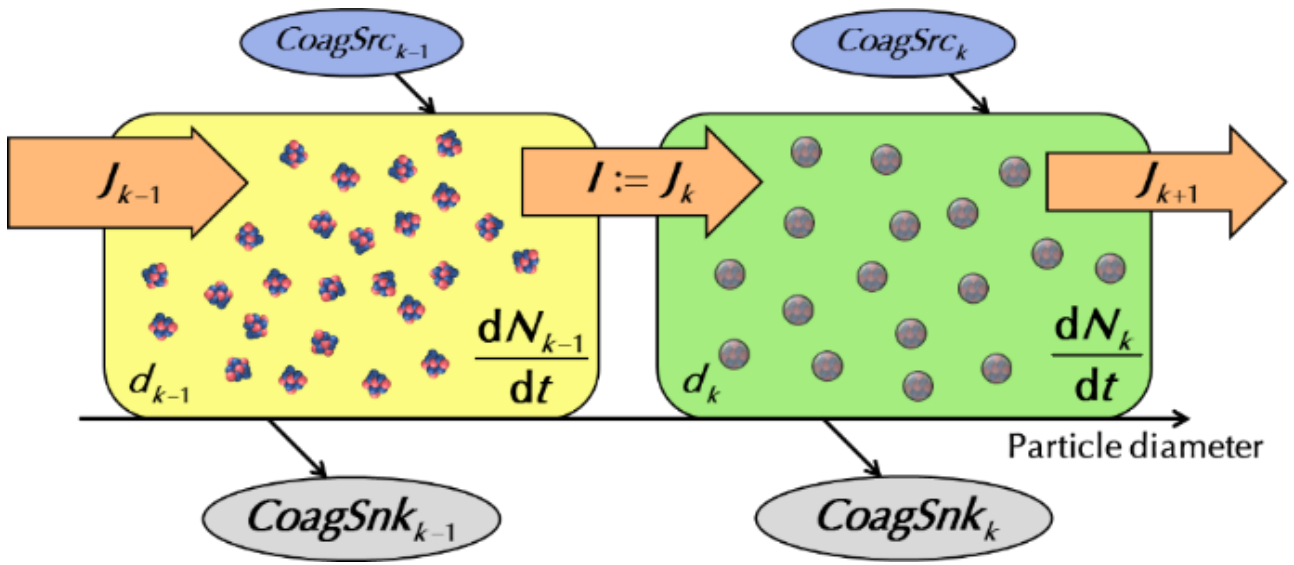


Figure 1: Schematic of the general dynamic equation.

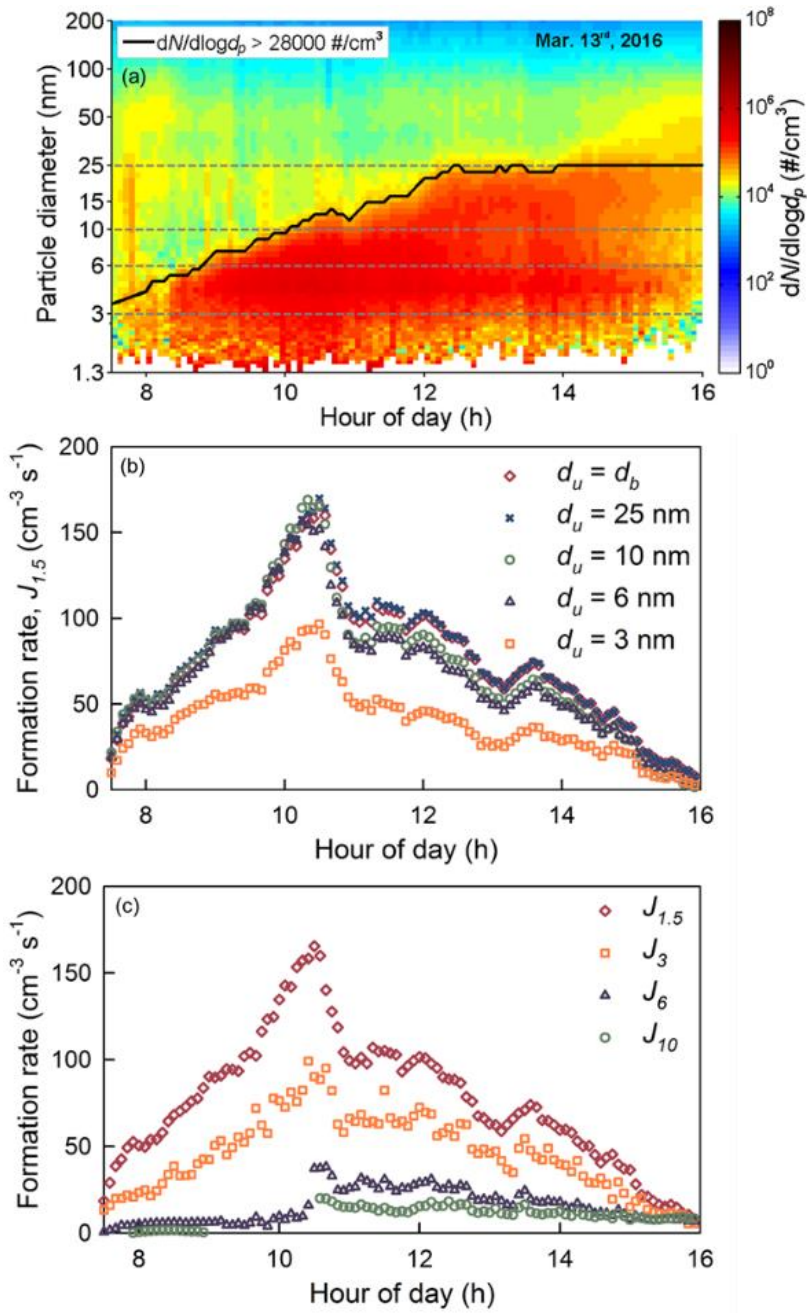


Figure 2: Comparison of formation rates estimated using different upper bounds, d_u . (a) A typical new particle formation event. Dashed gray lines represent different d_u in Eq. (1). Solid black lines corresponds to d_b , i.e., the varying upper bound determined by $dN/d\log d_p$. (b) Estimated formation rates with different upper bound, d_u , using Eq. (1). (c) Estimated formation rates with different d_k using Eq. (1). d_u equals 25 nm and d_{min} equals 1.3 nm in the four scatter plots.

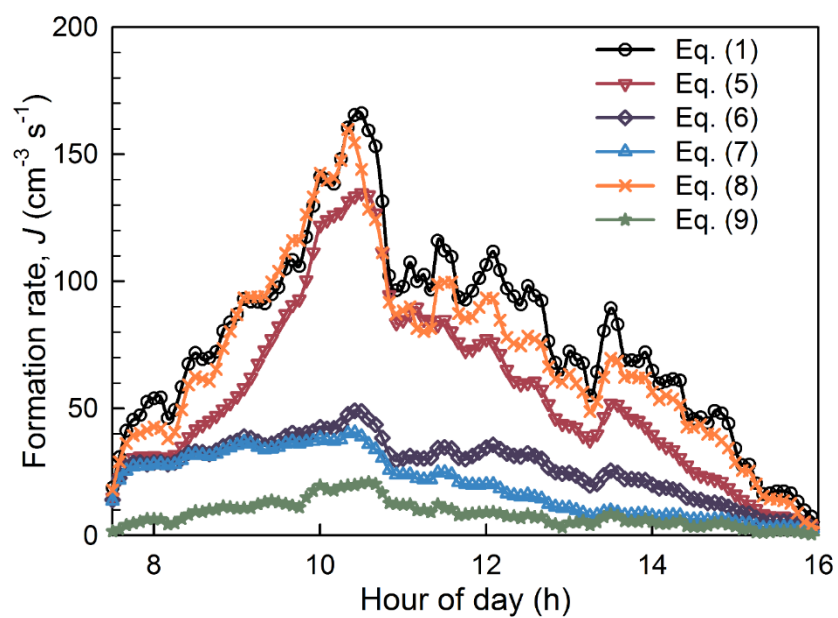


Figure 3: Comparison of formation rates estimated by different formulae.

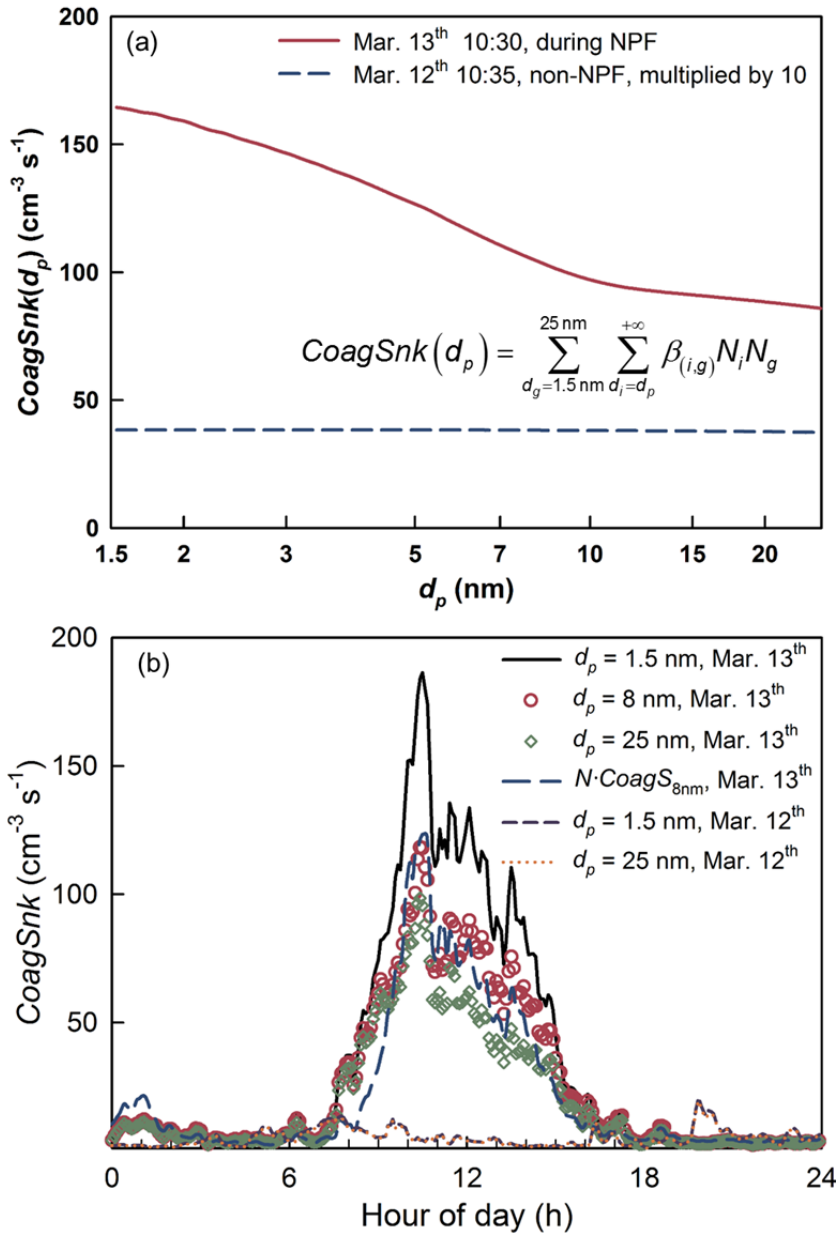


Figure 4: (a) $CoagSnk$ as a function of d_p , where d_p is the accounted minimum diameter when calculating $CoagS_g$ for particles at all different d_g , and scavenging due to coagulation with particles smaller than d_p is neglected, as the defined by the formula in panel (a). The dashed line corresponding to $CoagSnk$ on a non-NPF day is also monotonously decreasing with the increase of d_{min} by a negligible slope. (b) Time evolution of $CoagSnk$ versus time on a NPF day (Mar. 13th) and a non-NPF day (Mar. 12th). d_p is defined the same with that in panel (a). N is the number concentration of particles in the size range from 1.5 nm to 25 nm, while $CoagS_{8nm}$ is calculated using Eq. (3).

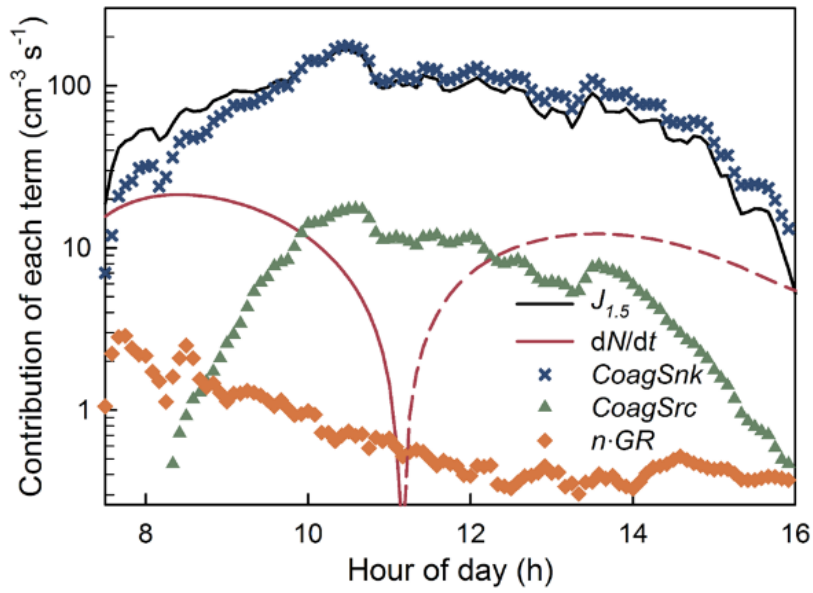


Figure 5: Contribution of each term to the estimated formation rate. dN/dt is obtained by fitting and shown in absolute value with solid and dashed lines corresponding to positive and negative parts, respectively. Note the upper bound, d_u , equals d_b as defined section 4.1 for better accuracy, however, it doesn't affect the generality of the result.

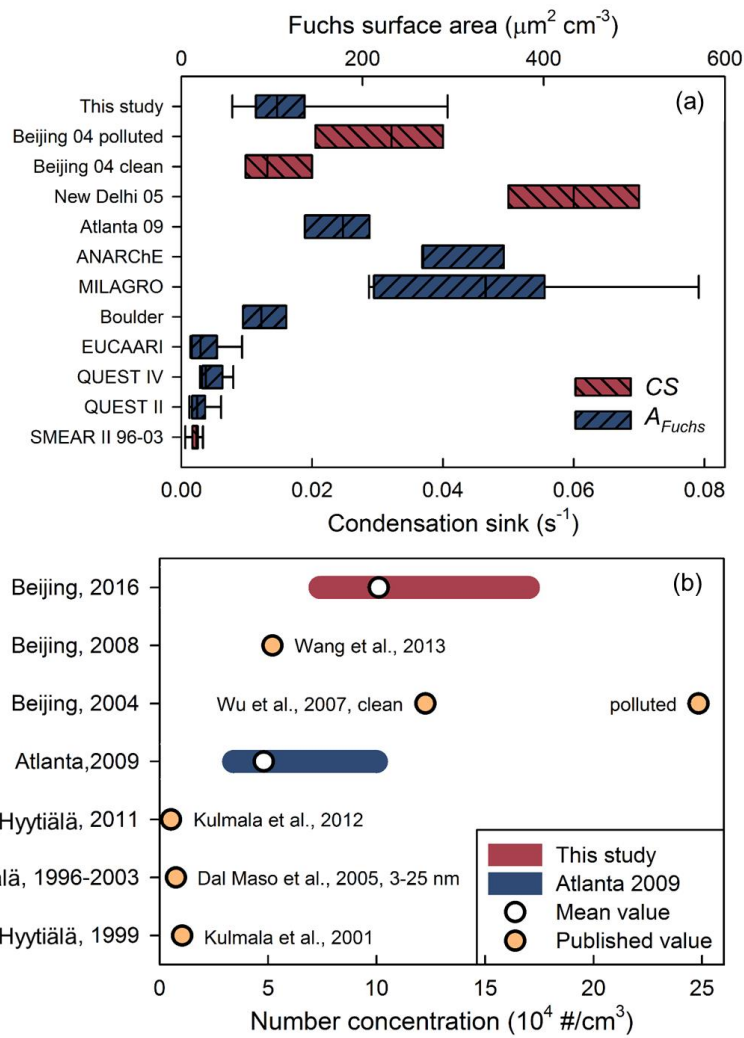


Figure 6: (a) Comparison of Fuchs surface area and condensation sink in Beijing (when NPF events occurred) with those in other locations. NPF days were classified by condensation sink in urban Beijing in 2004 (Wu et al., 2007). Condensation sink on NPF days in New Delhi was reported by Kulmala et al. (2005). ANARChE (McMurry et al., 2005) and MILAGRO (Iida et al., 2008) were conducted in Atlanta and Tecamac, respectively, while EUCCARI (Manninen et al., 2009), QUEST II (Sihto et al., 2006), QUEST IV (Riipinen et al., 2007) was conducted in SMEAR II (Dal Maso et al., 2005), Hyytiälä. A_{Fuchs} data in MILAGRO, ANARChE, Boulder, EUCCARI, QUEST II, and QUEST IV were published in Kuang et al. (2010). The ends of coloured rectangular correspond quartiles, while error bar represents the 10th and 90th percent value. (b) Comparison of peak number concentration of particles larger than 3 nm during NPF events in this study with those in Atlanta and other published data. Note that the published values (light orange points) in previous studies are not necessarily the mean values of the whole campaign periods.

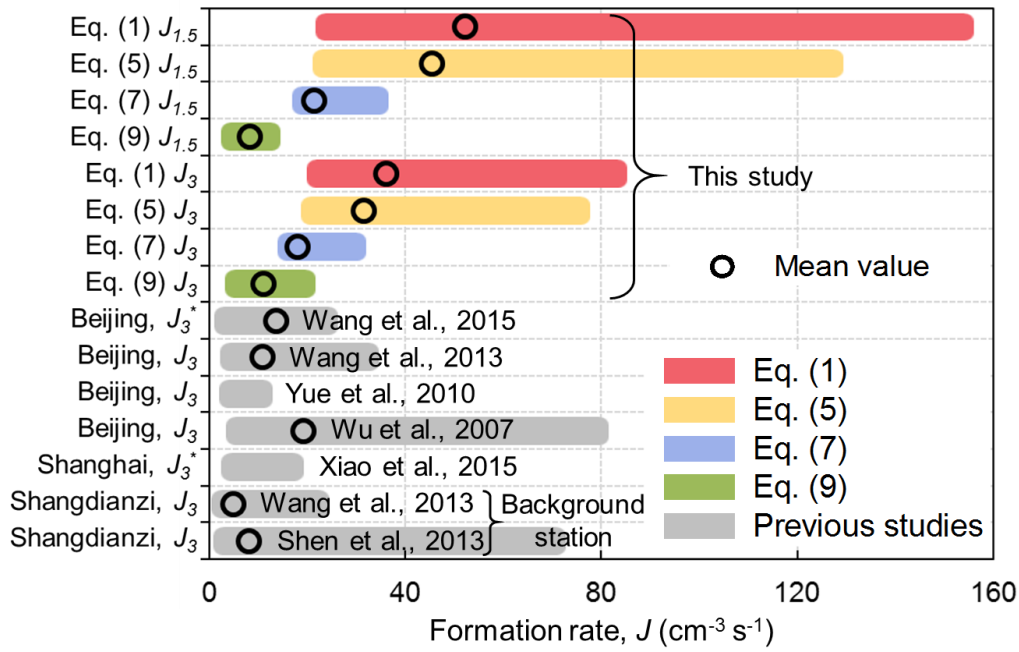
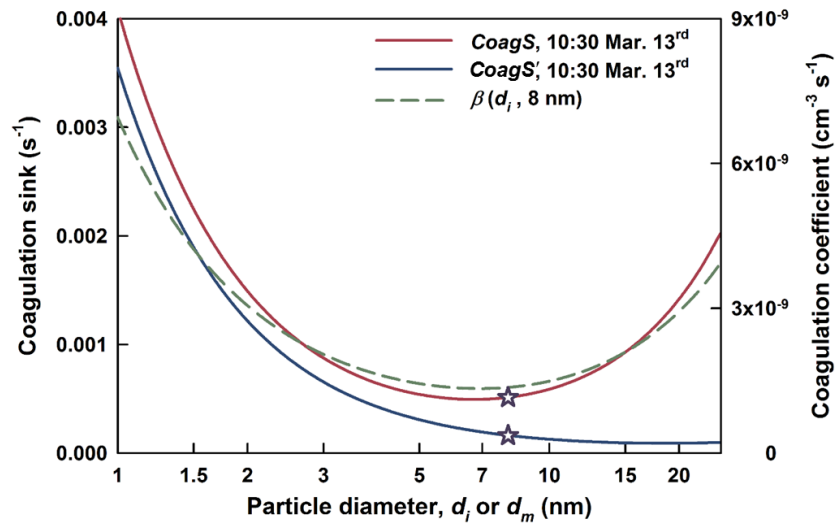


Figure 7: Estimated $J_{1.5}$ and J_3 using different equations. Previously reported J_3 in China were included for comparison. The ends of coloured rectangular correspond to the minimum value and the maximum value, respectively. *: The upper size bound to estimate formation rate, d_u , is 6 nm (rather than 25 nm) in Wang et al., 2015 and Xiao et al., 2015.



669

670 **Figure B1:** Coagulation coefficient and calculated coagulation sink during a typical NPF event. $CoagS$ and $CoagS'$ are defined in
671 Eq. (B7) and Eq. (B8), respectively, and d_m in this figure is treated as a variable rather than a constant value. The upper and lower
672 star denote $CoagS'_{8nm}$ and $CoagS_{8nm}$ which are used in the second term in the RHS of Eq. (5) and Eq. (6) to approximate $CoagS_{nk}$,
673 respectively.

674

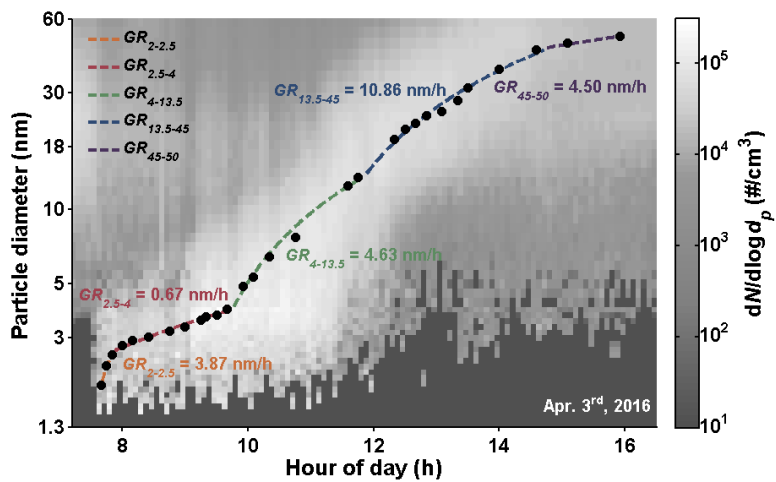


Figure B2: Size and time dependent growth rate on a NPF day observed in Beijing. Representative diameters are obtained by lognormal fitting of nucleation mode particles in each time bin, and GR is linearly fitted in each section.