



# 1 **Insight into winter haze formation mechanisms based on aerosol** 2 **hygroscopicity and effective density measurements**

3 Yuanyuan Xie, Xingnan Ye\*, Zhen Ma, Ye Tao, Ruyu Wang, Ci Zhang, Xin Yang, Jianmin Chen, Hong  
4 Chen

5 Shanghai Key Laboratory of Atmospheric Particle Pollution and Prevention (LAP<sup>3</sup>), Department of  
6 Environmental Science and Engineering, Fudan University, Shanghai 200433, China.

7 \*Correspondence to: Xingnan Ye (yexingnan@fudan.edu.cn), Jianmin Chen (jmchen@fudan.edu.cn).

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9 **Abstract:** We characterize a representative haze event from a series of periodic particulate matter (PM)  
10 episodes that occurred in Shanghai during winter 2014. Particle size distribution, hygroscopicity, and  
11 effective density were measured online, along with analysis of water-soluble inorganic ions and single  
12 particle mass spectrometry. Regardless of pollution level, the mass ratio of SNA/PM<sub>1.0</sub> (sulfate, nitrate,  
13 and ammonium) slightly fluctuated around 0.28 over the whole observation, suggesting that both  
14 secondary inorganic compounds and carbonaceous aerosols (including soot and organic matter)  
15 contributed substantially to the haze formation. Nitrate was the most abundant ionic species during hazy  
16 periods, indicating that NO<sub>x</sub> contributed more to haze formation in Shanghai than did SO<sub>2</sub>. The calculated  
17 PM concentration from particle size distribution displayed a variation pattern similar to that of measured



18 PM<sub>1.0</sub> during the representative PM episode, indicating that enhanced pollution level was attributable to  
19 the elevated number of larger particles. The number fraction of the near-hydrophobic group increased as  
20 the PM episode developed, indicating accumulation of local emissions. Three “banana-shape” particle  
21 evolutions were consistent with the rapid increase in PM<sub>1.0</sub> mass loading, indicating rapid size growth by  
22 condensation of condensable materials was responsible for the severe haze formation. Both  
23 hygroscopicity and effective density of the particles increased considerably with growing particle size  
24 during the banana-shaped evolutions, indicating that secondary transformation of NO<sub>x</sub> and SO<sub>2</sub> was a  
25 major contributor to the particle growth. Our results suggest that the accumulation of gas-phase and  
26 particulate pollutants under stagnant meteorological conditions and subsequent rapid particle growth by  
27 secondary processes, were primarily responsible for the haze pollution in Shanghai during wintertime.

28 **Keywords:** air pollution; size distribution; hygroscopic growth; secondary process; Shanghai.

29

## 30 1. Introduction

31 Atmospheric aerosol has significant influences on radiation balance and climate forcing of the  
32 atmosphere (Wang et al., 2011; Wang et al., 2014c; Wu et al., 2016a; IPCC, 2013); as well as strong impacts  
33 on visibility (Yang et al., 2012; Lin et al., 2014; Xiao et al., 2014) and public health (Heal et al., 2012) in  
34 heavily polluted areas. Recent studies found that short-term exposure to haze pollution could cause airway  
35 inflammation and aggravate respiratory symptoms in chronic obstructive pulmonary disease patients (Wu



36 et al., 2016b;Guan et al., 2016).

37 With the huge achievements in economic development and rapid urbanization over the past 30 years,  
38 particulate pollution has become a major environmental concern in China. The most severe haze event  
39 that occurred in the first quarter of 2013, spread over 1.6 million km<sup>2</sup> (Wang et al., 2014a). This event  
40 motivated the release of the Action Plan on Prevention and Control of Air Pollution with the goal of  
41 reducing PM<sub>2.5</sub> (particulate matter smaller than 2.5 μm in aerodynamic diameter) concentration by 15–25%  
42 in 2017 against 2012 in three major city clusters  
43 ([http://english.mep.gov.cn/News\\_service/infocus/201309/t20130924\\_260707.htm](http://english.mep.gov.cn/News_service/infocus/201309/t20130924_260707.htm)). In order to reduce  
44 the PM<sub>2.5</sub> concentration, extensive studies have been conducted to investigate the sources and formation  
45 mechanisms of haze pollution in recent years (Ye et al., 2011;Sun et al., 2016;Qiao et al., 2016;Hu et al.,  
46 2016;Li et al., 2016;Guo et al., 2014;Zheng et al., 2015;Guo et al., 2013;Wang et al., 2016;Peng et al.,  
47 2016). However, the haze formation mechanisms and source appointment of fine particles remain  
48 uncertain.

49 Guo et al. (2013) summarized historical reports from 2000 to 2008 in Beijing and found that the  
50 origins of urban fine particles varied in different seasons: the contribution of primary emissions is  
51 comparable to that of secondary formation during winter heating periods whereas secondarily produced  
52 aerosols dominate the fine PM sources in other seasons. As an important type of primary emissions in  
53 urban area, black carbon (BC) is primarily from incomplete fossil fuel combustion. Light absorption of



54 BC aerosols is increased after atmospheric aging by coating with secondary materials and restructuring  
55 (Khalizov et al., 2009). Due to cooling effect at the surface and warming effect aloft, the enhanced light  
56 absorption and scattering by aged BC particles stabilize the atmosphere, hindering vertical transport of  
57 gaseous and particulate pollutants (Wang et al., 2013). BC aging occurs much more efficiently in the  
58 presence of highly elevated gaseous aerosol precursors so that light absorption increases by a factor of  
59 2.4 within 4.6 h under highly polluted conditions in Beijing, significantly exacerbating pollution  
60 accumulation and strongly contributing to severe haze formation (Peng et al., 2016).

61 Due to the implement of several effective regulatory policies, the increasing trend of primary  
62 emissions has been under control since the 11<sup>th</sup> five-year period. A growing number of studies suggested  
63 that secondary production was the major contributor to the haze episodes in recent years (Shi et al.,  
64 2014;Zhao et al., 2013;Zhang et al., 2015a;Huang et al., 2014), in contrast that primary emissions were  
65 of great importance in some haze episodes (Niu et al., 2016). Guo et al. (2014) reported that the  
66 development of PM episodes in Beijing was characterized by efficient nucleation and continuous particle  
67 growth over an extend period dominated by local secondary formation. They attributed the continuous  
68 growth of particle size and constant accumulation of particle mass concentration to the highly elevated  
69 concentrations of gaseous precursors such as NO<sub>x</sub>, SO<sub>2</sub>, and volatile organic compounds (VOCs), while  
70 the contribution from primary emissions and regional transport was negligible. However, the role of  
71 regional transport of PM<sub>2.5</sub> in haze formation remains controversial (Li et al., 2015;Zhang et al., 2015b).



72 The most important advances in the understanding of urban PM formation were reviewed by Zhang  
73 et al. (2015c). The concentrations of SO<sub>2</sub>, NO<sub>x</sub>, and anthropogenic source VOCs in Beijing and other  
74 cities of the developing world are significantly higher than those in the urban areas of developed countries,  
75 resulting in large secondary production of sulfate, nitrate, and SOA. Synergetic effects among various  
76 organic and inorganic compounds may exist under highly polluted conditions, indicating different PM  
77 formation rates between developing and developed urban regions. Indeed, a large enhancement of  
78 particulate sulfate was typically observed during regional haze events in China (Chen et al., 2016; Wang  
79 et al., 2015; Fu et al., 2008; Xie et al., 2015). Currently, the highly elevated sulfate concentration during  
80 haze episodes cannot be fully explained by model simulations (Wang et al., 2014b; Chen et al., 2016).  
81 Recently, a significant breakthrough made by Wang et al. (2016) has provided a reasonable explanation  
82 in the high level of sulfate during haze episodes. It is revealed by their laboratory experiments that aqueous  
83 oxidation of SO<sub>2</sub> by NO<sub>2</sub> proceeds more efficiently with the increase of NO<sub>2</sub> concentration whereas the  
84 reaction is suppressed in acid conditions, because acid effect reduces the solubility of SO<sub>2</sub> and reaction  
85 rate. The enhanced sulfate formation during severe haze periods in Beijing was attributable to aqueous  
86 oxidation of SO<sub>2</sub> by NO<sub>2</sub> on hygroscopic fine particles under conditions of elevated RH and the  
87 concentrations of NH<sub>3</sub> and NO<sub>2</sub>, which was confirmed by the comparable SO<sub>2</sub> uptake coefficients for  
88 sulfate formation from field and laboratory results.

89 The hygroscopic properties of ambient particles vary significantly depending on the origin of the air



90 masses and the atmospheric aging process. In urban air, the population of near-hydrophobic particles can  
91 be assumed to consist largely of freshly emitted combustion particles containing high mass fractions of  
92 soot and water-insoluble organic compounds (Swietlicki et al., 2008; Massling et al., 2009). In contrast,  
93 secondary sulfate or nitrate aged particles are more-hygroscopic, and their relative abundance is primarily  
94 responsible for the hygroscopic growth of ambient particles at elevated RH (Topping et al.,  
95 2005; Aggarwal et al., 2007; Gysel et al., 2007). Thus, hygroscopicity can serve as a tracer of source origins,  
96 mixing state, and aging mechanisms of ambient particles. For example, temporal variations of aerosol  
97 hygroscopicity have helped the explanation of haze formation mechanisms in Beijing and Shanghai (Ye  
98 et al., 2011; Guo et al., 2014).

99 Density is one of the most important physicochemical properties for atmospheric aerosols. Effective  
100 density has served as a tracer for new particle formation and for the aging process in previous studies (Yin  
101 et al., 2015; Guo et al., 2014). The ambient particles in urban areas are mostly complex mixtures of  
102 elemental carbon (EC), organics (OC), and secondary inorganic aerosols (SIA) (Hu et al., 2012). The  
103 effective densities of traffic particles are below  $1.0 \text{ g cm}^{-3}$ , and density decreases with the increase of  
104 particle size because there are more voids between primary particles in relatively larger aggregates  
105 (Momenimovahed and Olfert, 2015). The density of OC is in between those of EC and SIA, and varies  
106 with source. The effective density of combustion particles increases by filling the voids in the agglomerate  
107 particles with condensed semi-volatile materials, or by restructuring agglomerates with hygroscopic SIA



108 (Momenimovahed and Olfert, 2015;Zhang et al., 2008).

109 In this study, a combined HTDMA-APM system was used to investigate the variations of  
110 hygroscopicity and effective density of submicrometer aerosols during winter 2014 in urban Shanghai. In  
111 addition, cascade samples were collected and a single particle mass spectrometry was used to better  
112 understand the hygroscopicity and density variations. The primary objectives of this study were to  
113 investigate the particle growth mechanisms and to identify the contribution of local emissions during the  
114 winter haze episode.

115

## 116 **2. Experimental**

### 117 **2.1. Sampling site**

118 The measurements were conducted from December 21, 2014 to January 13, 2015 at the Department of  
119 Environmental Science and Engineering in the main campus of Fudan University (31.30°N, 121.5°E), a  
120 representative urban site close to a sub-center of Shanghai (Ye et al., 2010). At a supersite about 100 m  
121 away, PM<sub>1.0</sub> was monitored using a Thermo Scientific™ 5030 SHARP monitor. Trace gas pollutants were  
122 monitored using Thermo Scientific™ i-series gas analyzers (43i for SO<sub>2</sub>, 49i for O<sub>3</sub>, 42i for NO/NO<sub>2</sub>/NO<sub>x</sub>),  
123 and meteorological data were monitored using an automatic meteorological station (Model CAWS600,  
124 Huayun Inc., China) (Yin et al., 2015). The concentrations of PM<sub>2.5</sub>, PM<sub>10</sub>, and CO were released by the  
125 Shanghai Environmental Monitoring Center. The height of the Planet Boundary Layer (PBL) was



126 computed online using the NCEP Global Data Assimilation System (GDAS) model  
127 (<http://ready.arl.noaa.gov/READYamet.php>).

## 128 **2.2. HTDMA-APM system**

129 Particle size distribution, hygroscopic growth factor (GF), and effective density were measured using  
130 a custom-built HTDMA-APM system (Figure 1). The custom-built HTDMA (Hygroscopic Tandem  
131 Differential Mobility Analyzers) mainly consist of two long DMAs (3081L, TSI Inc.), a humidifier (PD-  
132 50T-12MSS, Perma Pure Inc.) and a Condensation Particle Counter (CPC, Model 3771, TSI Inc.). A  
133 detailed description of the HTDMA is available in Ye et al. (2009). In this observation, particle number  
134 size distribution in the range of 14–600 nm and hygroscopic growth at 83% RH for particles with dry  
135 diameters of 40, 100, 220, 300, 350, and 400 nm were determined by HTDMA (in turn). The  
136 determination of effective density by DMA-APM was described previously (Yin et al., 2015; Pagels et al.,  
137 2009). Briefly, a combined system consisting of a compact Aerosol Particle Mass Analyzer (APM, Model  
138 3601, Kanomax Inc.) and a CPC (Model 3775, TSI Inc.) was connected to the sample tubing through a  
139 3-way electrical switch behind the upstream DMA (DMA1). The APM comprises two coaxial cylindrical  
140 electrodes rotating at the same angular velocity. Charged aerosol particles of a certain diameter sized by  
141 DMA1 are axially fed into the annular gap between the electrodes and experienced an outward centrifugal  
142 force from the particle rotating and an inward electrostatic force from the high-voltage field between the  
143 electrodes. Particles pass through the APM and are sent to the CPC when the two forces are balanced.





144 The mass of particles that pass through the APM is determined by the rotation rate and the applied voltage.  
145 Effective densities for dry diameters of 40, 100, 220, and 300 nm were determined by the method of  
146 DMA-APM in this study. The HTDMA-APM was operated alternatively in HTDMA mode and then  
147 DMA-APM mode, for every 40 min.

148 Before the field observation, the HTDMA-APM was calibrated using 40–450 nm NIST-Traceable PSL  
149 particles and ammonium sulfate. The measured HTDMA data were inversed with the TDMA<sub>inv</sub> algorithm  
150 to obtain the actual GF distribution. This is because the raw data are only a skewed and smoothed  
151 integral transform of the actual growth factor probability density function (GF-PDF) (Gysel et al., 2009).  
152 The hygroscopicity parameter  $\kappa$  was derived from the GF data after inversion with the TDMA<sub>inv</sub> algorithm  
153 according to the  $\kappa$ -Köhler theory (Petters and Kreidenweis, 2007).

### 154 **2.3. SPAMS**

155 A Single Particle Aerosol Mass Spectrometry (SPAMS, Hexin Analytical Instrument Co., Ltd., China)  
156 installed in the same room with the HTDMA-APM system was used to obtain the chemical and size  
157 information of individual particles in the range of 0.2-2  $\mu\text{m}$ . Detailed information on SPAMS is available  
158 in Li et al. (2011). Briefly, ambient particles are drawn into a vacuum chamber through an aerodynamic  
159 focusing lens and accelerated to a size-dependent terminal velocity. Sized particles are desorbed and  
160 ionized by the pulsed desorption/ionization laser (Q-switched Nd: YAG,  $\lambda=266$  nm) at the ion source  
161 region. Both positive and negative mass spectra for a single particle are recorded by a bipolar time-of-



162 flight spectrometer. The single particle information was imported into YAADA (version 2.11,  
163 www.yaada.org). Based on the similarities of the mass-to-charge ratio and peak intensity, particles were  
164 classified using the ART-2a method.

## 165 **2.4. Ion chromatography**

166 Cascade aerosol samples for offline analysis were collected at the roof platform of the Environmental  
167 Building using a 10-stage MOUDI sampler (Micro-Orifice Uniform Deposit Impactor, Model 110-NR,  
168 MSP Corp., USA). Detailed description of the sampling, pretreatment, chemical analysis, and quality  
169 control of this system is available in Tao et al. (2016). Briefly, cascade samples were collected every 24  
170 h using the PALL7204 quartz filter as the collection substrate. Each filter was weighted with a BP211D  
171 electronic balance at  $25\pm 1^\circ\text{C}$  and  $40\pm 2\%\text{RH}$ . The water extract of each sample was analyzed using an Ion  
172 Chromatograph (Metrohm 883 basic IC plus, Switzerland) equipped with a third-party column heater  
173 (CT-100, Agela Corp., China). Seven anions ( $\text{F}^-$ ,  $\text{Cl}^-$ ,  $\text{NO}_2^-$ ,  $\text{Br}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$  and  $\text{PO}_4^{3-}$ ) were resolved  
174 using a Metrosep A Supp 5-250/4.0 column at  $35^\circ\text{C}$  with an eluent of  $3.2\text{ mmol L}^{-1}\text{ Na}_2\text{CO}_3 + 1.0\text{ mmol}$   
175  $\text{L}^{-1}\text{ NaHCO}_3$ . Six cations ( $\text{Li}^+$ ,  $\text{Na}^+$ ,  $\text{NH}_4^+$ ,  $\text{K}^+$ ,  $\text{Ca}^{2+}$ , and  $\text{Mg}^{2+}$ ) were separated by a Metrosep C4-250/4.0  
176 column at  $30^\circ\text{C}$  with an eluent of  $1.7\text{ mmol L}^{-1}\text{ HNO}_3 + 0.7\text{ mmol L}^{-1}$  2,6-pyridine dicarboxylic acid.

177

## 178 **3. Results and discussion**

### 179 **3.1. Periodic cycle of PM episodes during the observation period**



180 Figure 2 shows the temporal variations of PM mass loading during the winter observation (December  
181 21, 2014 to January 13, 2015). The official data of PM<sub>2.5</sub> and PM<sub>10</sub> were blank on some clean days.  
182 Meteorologically, our measurement was deployed in a typical winter period. The average concentrations  
183 of PM<sub>1.0</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub> were  $57.3\pm37.0$ ,  $87.2\pm67.2$ , and  $127.8\pm77.7$   $\mu\text{g m}^{-3}$ , respectively. About 62%  
184 of hourly averaged PM<sub>2.5</sub> concentrations exceeded  $75$   $\mu\text{g m}^{-3}$  of the Chinese Grade II guideline (GB 3095-  
185 2012), indicating heavy particle pollution in Shanghai during wintertime. The PM episodes exhibited a  
186 clear periodic cycle of  $\sim 5$  days. A similar feature was previously observed in Beijing (Guo et al., 2014).  
187 At the beginning of each cycle, the PM<sub>1.0</sub> level was below  $35$   $\mu\text{g m}^{-3}$ . During the clean period, the  
188 differences among the concentrations of PM<sub>1.0</sub>, PM<sub>2.5</sub>, and PM<sub>10</sub> were insignificant. Occasionally the  
189 measured PM<sub>2.5</sub> concentrations were larger than those of PM<sub>10</sub>, possibly due to system error. However,  
190 the particle mass concentration began to increase in the next few days, with PM<sub>1.0</sub> and PM<sub>2.5</sub> peaking at  
191 over  $100$  and  $200$   $\mu\text{g m}^{-3}$ , respectively. During the late episodes, the PM mass loading abruptly dropped,  
192 due to change in the atmospheric dilution or wet deposition.

### 193 **3.2 Contributions of secondary inorganic aerosols to PM<sub>1.0</sub> mass loading**

194 Figure 3 illustrates the daily concentrations of sulfate, nitrate, and ammonium as a function of PM<sub>1.0</sub>  
195 mass loading. In general, the sum of concentrations of sulfate, nitrate, and ammonium (SNA) increased  
196 linearly as PM<sub>1.0</sub> mass loading increased. It is noticeable that the SNA/PM<sub>1.0</sub> ratio slightly fluctuated  
197 around 0.28, regardless of the pollution level. Because soil dust and sea salt made a negligible contribution



198 to the fine particle mass concentration in this study, the almost constant ratio of SNA/PM<sub>1.0</sub> indicates that  
199 SNA and carbonaceous aerosols (including soot and organic matter) synchronously increased during haze  
200 episodes. As the PM<sub>1.0</sub> concentration increased, the concentration of nitrate increased more rapidly than  
201 sulfate so that it became the most abundant ionic species at PM<sub>1.0</sub> > 40 μg m<sup>-3</sup>. This indicated that NO<sub>x</sub>  
202 contributed more to haze formation in Shanghai than did coal-fired sources. Generally, the visibility  
203 decreased with the increase in PM concentration, indicating photochemical activity began to weaken as  
204 the development of haze episodes. The large increase in nitrate concentration may be attributable to  
205 heterogeneous reaction on the preexisting particles. Nitrate formation is highly dependent on the surface  
206 area of preexisting particles and is favored under NH<sub>3</sub>-rich conditions (Chu et al., 2016). In contrast, Han  
207 et al. (2016) reported that the mass ratio of nitrate to sulfate decreased with increase of PM<sub>2.5</sub> level and  
208 that the sources of sulfate contributed more to haze formation in Beijing than mobile sources. This finding  
209 suggests that the haze formation mechanism is likely different in Shanghai and Beijing. VOCs and NO<sub>x</sub>  
210 are exclusively from local emissions whereas regional transport is a big source of SO<sub>2</sub> under stagnant  
211 atmosphere, due to their different atmospheric lifetimes (Guo et al., 2014). Considering the relatively  
212 smaller contribution of sulfate, our results reveal that the accumulation and secondary transformation of  
213 local emissions likely played a dominant role in this haze formation.

### 214 3.3 Aerosol hygroscopicity and effective density during the observation period

215 Figure 4a displays a box chart of the median hygroscopicity of each hygroscopic growth factor



216 distribution for different sizes. Considering all of the growth factor distributions collectively, the  
217 hygroscopicity parameter  $\kappa$  increased with increase of the dry diameter, with an average  $\kappa$  of 0.161 at 40  
218 nm and 0.338 at 300 nm. Assuming a two-component system of a model salt (ammonium sulfate,  $\kappa_m =$   
219 0.53) and an insoluble species ( $\kappa = 0$ ), the volume fraction of hygroscopic species ( $\varepsilon$ ) can be obtained  
220 based on the Zdanovsldi-Stokes-Robinson (ZSR) mixing rule. The average  $\varepsilon$  was 0.3 for 40 nm particles,  
221 suggesting that the primary particles or initial growth of freshly generated particles were dominated by  
222 non-hygroscopic species. In contrast, the 300 nm particles were extremely aged, with more-hygroscopic  
223 species.

224 Generally, the HTDMA-measured hygroscopicity was limited to the size range below 250 nm, and it  
225 is common that the GF increases with increase of particle size. The increase of aerosol hygroscopicity  
226 with size was attributed to the addition of more-hygroscopic SNA (Swietlicki et al., 2008; Ye et al., 2010).  
227 Gasparini et al. (2006) reported that the GF first increased and then decreased with increase of particle  
228 size, peaking at the diameter of 300 nm. In this study, the determination size range was extended to 400  
229 nm and no decrease in GF was observed. We attribute the different hygroscopicity to the large emissions  
230 of  $\text{SO}_2$  and  $\text{NO}_x$  in China, which were responsible for the strong formation of sulfate and nitrate. The  
231 variation of hygroscopicity parameter  $\kappa$  was much greater for 40 nm particles. The particle population  
232 with  $\kappa < 0.1$  was attributed to fresh traffic particles (Ye et al., 2013). The considerable percentile of  $\kappa <$   
233 0.1 indicated that the 40 nm particle population was sometimes dominated by near-hydrophobic particles.



234 Figure 4b displays a box chart of median effective density for different particle sizes. The median  
235 effective density varied in the narrow range of  $\rho_{\text{eff}} = 1.35\text{--}1.41 \text{ g cm}^{-3}$  for 40–300 nm particle population.  
236 The size distribution of particle density varied in the literature. Hu et al. (2012) and Yin et al. (2015)  
237 reported that particle density increased as particle size increased while a contradictory trend was observed  
238 by Geller et al. (2006) and Spencer et al. (2007). The difference was attributable to the contribution of  
239 fresh traffic particles. Although the dominant accumulation mode particles have an effective density  
240 greater than Aitken mode ones, the presence of a lower effective density group associated with emissions  
241 from traffic exhaust might decrease the mean effective density to a value lower than that of Aitken mode  
242 particles (Levy et al., 2014). Yin et al. (2015) reported that a quasi-monodisperse density distribution was  
243 dominant for accumulation mode particles. In contrast, externally mixed aerosols with a lower density  
244 group ( $\rho_{\text{eff}} = \sim 1.0 \text{ g cm}^{-3}$ ) were often present in this observation, and were responsible for the decrease of  
245 the mean effective density. The lower effective density group was attributed to fresh or slightly aged  
246 traffic-related particles, because the number fraction of the lower density group increased as the  
247 concentration of NO increased.

### 248 **3.4 Characteristics of a representative PM episode**

249 The PM episode from January 7 to 12 was a representative case of severe haze formation and  
250 elimination processes. It can be divided into clean (January 7), transition (January 8), haze (January 9–11),  
251 and post-haze (January 12) periods. During the transition from the clean to haze period (January 7 to 8),



252 both  $PM_{1.0}$  and  $PM_{2.5}$  concentrations increased slightly, with an average  $PM_{1.0}/PM_{2.5}$  ratio of 0.65. A sharp  
253 increase in  $PM_{2.5}$  (of  $125 \mu\text{g m}^{-3}$ ) was observed from 6:00 to 12:00 local time on the morning of January  
254 9. During the haze period, the concentration of  $PM_{2.5}$  exceeded  $115 \mu\text{g m}^{-3}$  (medially polluted level,  
255 HJ633-2012) for 63 h. On January 11, the hourly  $PM_{2.5}$  concentration exceeded  $250 \mu\text{g m}^{-3}$ , corresponding  
256 to the severely polluted level.

257 Figure 5 displays the temporal evolution of particle size distribution in comparison with the measured  
258  $PM_{1.0}$  concentration during the representative PM episode. The calculated PM concentrations ( $PM_{\text{cal}}$ )  
259 were obtained based on the particle size distribution and average effective density of  $1.39 \text{ g m}^{-3}$  in the  
260 range of 14–600 nm measured in this study. In contrast to the fact that particle size distribution was  
261 dominated by nanoparticles during the clean period, the burst of Aitken mode particles and subsequent  
262 continuous growth to approximately 200 nm in diameter was observed three times during the haze period,  
263 indicating that the presence of numerous larger particles is likely responsible for the severe particle  
264 pollution (Guo et al., 2014). The importance of larger particles in haze formation is also illustrated by the  
265 contour plot of the particle volume size distribution. The difference of total number concentration between  
266 transition and haze periods was insignificant, whereas the volume concentration increased rapidly during  
267 the haze period. This feature clearly demonstrates that the haze formation was closely correlated with  
268 particle growth and elevated number of larger particles.

269 Interestingly, the particle mass concentration was sensitive to variations of wind speed and planetary



270 boundary layer (PBL). During the transition and haze periods, the wind speed decreased considerably  
271 with insignificant change in prevailing wind (Figure S1). This indicated that outside transportation  
272 became less and less significant. It is noteworthy that the temporal evolution of the particle mass  
273 concentration was inversely correlated with the PBL height. The decreasing PBL provided a stagnant  
274 atmosphere that favored the accumulation of local emissions. This finding reveals that the severe haze  
275 pollution was likely triggered by the adverse meteorological conditions. The impact of decreasing PBL  
276 height on haze formation can also be evidenced by the variations of trace gaseous species (Figure S2).  
277 During the PM episode, the concentrations of NO<sub>2</sub>, SO<sub>2</sub>, and CO displayed variation trends similar to that  
278 of the particle concentration. The fluctuations of trace gas concentrations were caused by primary  
279 emission and secondary processes. Noticeably, the concentration of NO increased dramatically in rush  
280 hours during the haze period, whereas it fluctuated slightly during the clean period; indicating that local  
281 emissions were easily accumulated under stagnant atmosphere. In addition, the maximum concentration  
282 of O<sub>3</sub> remained considerably higher during daytime, whereas it decreased significantly at night. The most  
283 plausible explanation is that O<sub>3</sub> was consumed rapidly by the accumulating trace gases, such as NO<sub>x</sub>, and  
284 VOCs.

### 285 **3.5 Variations of hygroscopicity and effective density during the PM episode**

286 Figure 6 shows the averaged hygroscopicity and effective density for different pollution periods of the  
287 PM episode. Regardless of the pollution period, near- hydrophobic particles were externally mixed with





288 some hygroscopic particles. During the clean period, the more-hygroscopic particles dominated the 40  
289 nm particle population, indicating that the near-hydrophobic primary particles were rapidly dispersed due  
290 to atmospheric dilution. The number fraction of the near-hydrophobic group for different sizes increased  
291 as the PM episode developed, indicative of the increasing accumulation of local emissions. Notably, the  
292 increase of the near-hydrophobic particles with the evolution of the PM episode become less significant  
293 as particle size increased, indicating that primary emission exerted a more significant impact on smaller  
294 particles than on larger ones. The median diameter of nascent traffic particles from various gasoline  
295 sources ranged between 55 and 73 nm with an average of 65 nm (Momenimovahed and Olfert, 2015).  
296 Therefore, the number fraction of the near-hydrophobic particles larger than 200 nm is not sensitive to  
297 the accumulation of traffic emissions.

298 Interestingly, the variations of particle effective density for different sizes are in good agreement with  
299 the hygroscopicity. The dominant peak of effective density distribution appeared at  $\rho_{\text{eff}} = \sim 1.5 \text{ g cm}^{-3}$  for  
300 40 nm particles in the clean period, indicating that they are highly aged with hygroscopic inorganic salts  
301 (Yin et al., 2015). As the episode developed, the mean density shifted to lower values, indicating the  
302 increasing contribution of less-massive carbonaceous materials. The averaged density distribution was  
303 broadened as the episode developed, suggesting that it could be deconvolved into two groups and that the  
304 number fraction of the low-density group increased. This finding revealed that the less-massive particles  
305 are less hygroscopic whereas the larger density group corresponds to the more-hygroscopic one. In



306 addition, the variations of hygroscopicity and effective density coincided with the evolution of PBL height,  
307 indicating that the increasing accumulation of local emissions due to adverse atmospheric conditions is  
308 likely responsible for the enhancement of those near-hydrophobic and less-massive particles.

309 Figure 7 displays the temporal profiles for contributions of EC (including bare EC and OC-coated EC),  
310 OC, sulfate, and nitrate determined by SPAMS. Obviously, the relative contribution of nitrate increased  
311 as the episode developed. In contrast, the relative contribution of sulfate displayed an opposite trend. This  
312 feature is comparable with the aforementioned results of SNA, thus further highlighting the important  
313 role of nitrate in haze formation in Shanghai. The number fraction of EC particles generally increased  
314 during the haze period, peaking at midnight on January 9 and 10. It should be pointed out that the  
315 measured number fraction possibly underestimated the contribution of EC particles because the dominant  
316 size range of fresh traffic particles is below the detection limit of SPAMS (0.2–2.0  $\mu\text{m}$ ). This finding  
317 provides good support for the increase of near-hydrophobic and less-dense particles as the episode  
318 developed. Niu et al. (2016) reported that the number ratio of secondary particles to soot in haze samples  
319 was higher than that collected in the clean days in Beijing. Our finding is comparable to their results. In  
320 contrast, the number fraction of pure OC decreased during the pollution event. The possible explanation  
321 is that the condensation of organic matter was favored on the large amount of preexisting EC particles, or  
322 that photo-oxidation of VOCs was minimized due to lower solar radiation.

### 323 **3.6 Evolutions of hygroscopicity and effective density with particle growth**



324 Three “banana-shape” evolutions of the particle size distribution were identified in the representative  
325 PM episode. The banana-type contour plot of particle size distributions is a typical characteristics of new  
326 particle formation (NPF) events and traditionally regarded as one of the most important criteria for  
327 identifying NPF (Xiao et al., 2015; Dal Maso et al., 2005; Levy et al., 2013; Zhang et al., 2012).  
328 Atmospheric NPF is often defined by the burst of nucleation mode particles and subsequent growth of the  
329 nuclei to larger particles (Zhang et al., 2012; Kulmala et al., 2012). Gas-phase sulfuric acid produced via  
330 oxidation of SO<sub>2</sub> by OH radical plays a dominant role in the NPF events. NPF is typically completely  
331 suppressed when preexisting particles is abundant, because gas-phase sulfuric acid is rapidly lost to the  
332 surfaces of preexisting aerosols (Zhang et al., 2012). In addition to sulfuric acid, low-volatility organic  
333 species, and interaction between sulfate and organics are important for NPF (Zhang et al., 2004; Zhao et  
334 al., 2009). However, the possibility of NPF can be ignored in this study due to the absence of the burst of  
335 nucleation mode particles and the high concentration of PM<sub>1.0</sub>. The burst of Aitken mode particles was  
336 attributable to rapid accumulation of traffic emissions during rush hours under stagnant atmospheric  
337 conditions. The “banana-shape” evolutions were primarily caused by coagulation and condensation  
338 growth, which provided an excellent opportunity to reveal the chemical mechanism of particle growth  
339 during the PM episode.

340 The first “banana-shape” evolution of the particle size distribution occurred from approximately 05:00  
341 to 15:00 on January 9, with increase of the particle number concentration ( $N_{\text{total}}$ ) from  $1.7 \times 10^4$  to  $3.4 \times 10^4$



342  $\text{cm}^{-3}$  followed by a decrease trend until 17:00 (Period 1). The second “banana-shape” evolution occurred  
343 from approximately 18:00 on January 9 to approximately 12:00 on January 10 (Period 2). The  $N_{\text{total}}$   
344 increased from  $2.1 \times 10^4$  to  $4.2 \times 10^4 \text{ cm}^{-3}$  within 3 h, followed by gradual decrease of  $N_{\text{total}}$  in contrast to  
345 continuous increase of the particle mass concentration. During the growth process, the mode diameter of  
346 the particle population increased from below 40 nm to approximately 200 nm. The third “banana-shape”  
347 evolution began in the evening rush hours on January 10, with continuous increase of PM mass  
348 concentration for 12 h (Period 3).

349 Figure 8 illustrates the evolution of particle hygroscopicity and effective density during periods 2 and  
350 3. During the initial stage, the measured GF and effective density distributions were both bimodal, with  
351 a dominant peak at  $\text{GF} = \sim 1.0$  and  $\rho_{\text{eff}} = \sim 1.0 \text{ g cm}^{-3}$ , respectively. In a previous study, we found that the  
352 number fraction of near-hydrophobic particles varied with the traffic exhaust (Ye et al., 2013). Moreover,  
353 laboratory studies showed that the effective density of 50 nm vehicle particles was approximately  $1.0 \text{ g}$   
354  $\text{cm}^{-3}$  (Olfert et al., 2007; Park et al., 2003; Momenimovahed and Olfert, 2015). These findings indicate that  
355 the initial burst of Aitken mode particles is attributable to the presence of enhanced traffic-related  
356 emissions. In contrast, the number fraction and GF of the more-hygroscopic group increased with growing  
357 particle size, indicating the addition of hygroscopic inorganic species. The variation of the effective  
358 density of the particles was similar to that of the hygroscopicity, indicating the increase of high density  
359 materials. In general, inorganic sulfate and nitrate are more hygroscopic and denser than soot particles or



360 organic aerosols (Yin et al., 2015). These findings suggest that secondary sulfate and nitrate increased  
361 with the growing particle size, indicating the importance of the conversion of SO<sub>2</sub> and NO<sub>x</sub> in particle  
362 growth. This conclusion is supported by the largest SNA concentration in PM<sub>1.0</sub> during the PM episode  
363 (31.3 μg m<sup>-3</sup> on January 10 and 23.8 μg m<sup>-3</sup> on January 11). Considering that the concentration of nitrate  
364 was much higher than that of sulfate during the haze event, the increase of hygroscopicity was dominated  
365 by the addition of nitrate.

366

#### 367 4. Conclusions

368 Particle size distribution, size-resolved hygroscopic growth and effective density of sub-micrometer  
369 aerosols were determined using a HTDMA-APM system along with measurements of cascade samples  
370 and single particle mass spectrometry in urban Shanghai during winter 2014.

371 The PM episode exhibited a periodic cycle of ~5 days. The average concentration of PM<sub>2.5</sub> was  
372 87.2±67.2 μg m<sup>-3</sup>, with approximately 62% of hourly PM<sub>2.5</sub> concentrations exceeding the Chinese Grade  
373 II guideline. Both secondary inorganic salts and carbonaceous aerosols contributed substantially to haze  
374 formation, because the SNA/PM<sub>1.0</sub> ratio was almost constant during the observation period. Nitrate  
375 became the most abundant ionic species at PM<sub>1.0</sub> >40 μg m<sup>-3</sup>, indicating that the sources of nitrate  
376 contributed more to haze formation in Shanghai than did SO<sub>2</sub>.

377 The severe haze pollution was likely triggered by the adverse meteorological conditions, which caused



378 a large accumulation of local emissions and subsequent rapid growth to larger particles. As the PM  
379 episode developed, the number fraction of near-hydrophobic particles of different size increased,  
380 consistent with decrease of the mean effective density. Both hygroscopicity and effective density of the  
381 particles were found to increase considerably with growing particle size, indicating that secondary aerosol  
382 formation was a major contributor to particle growth. Our results suggest that the accumulation of local  
383 emissions under adverse meteorological conditions and subsequent rapid particle growth by secondary  
384 processes are primarily responsible for the haze pollution in Shanghai during wintertime.

385

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389

### 390 **Reference**

391 Aggarwal, S. G., Mochida, M., Kitamori, Y., and Kawamura, K.: Chemical closure study on hygroscopic  
392 properties of urban aerosol particles in Sapporo, Japan, *Environmental Science & Technology*, 41, 6920-  
393 6925, 10.1021/es063092m, 2007.

394 Chen, D., Liu, Z., Fast, J., and Ban, J.: Simulations of sulfate-nitrate-ammonium (SNA) aerosols during  
395 the extreme haze events over northern China in October 2014, *Atmospheric Chemistry and Physics*, 16,



- 396 10707-10724, 10.5194/acp-16-10707-2016, 2016.
- 397 Chu, B., Zhang, X., Liu, Y., He, H., Sun, Y., Jiang, J., Li, J., and Hao, J.: Synergetic formation of secondary  
398 inorganic and organic aerosol: effect of SO<sub>2</sub> and NH<sub>3</sub> on particle formation and growth, Atmospheric  
399 Chemistry and Physics, 16, 14219-14230, 10.5194/acp-16-14219-2016, 2016.
- 400 Dal Maso, M., Kulmala, M., Riipinen, I., Wagner, R., Hussein, T., Aalto, P. P., and Lehtinen, K. E. J.:  
401 Formation and growth of fresh atmospheric aerosols: eight years of aerosol size distribution data from  
402 SMEAR II, Hyytiälä, Finland, Boreal Environment Research, 10, 323-336, 2005.
- 403 Fu, Q. Y., Zhuang, G. S., Wang, J., Xu, C., Huang, K., Li, J., Hou, B., Lu, T., and Streets, D. G.:  
404 Mechanism of formation of the heaviest pollution episode ever recorded in the Yangtze River Delta, China,  
405 Atmospheric Environment, 42, 2023-2036, 2008.
- 406 Gasparini, R., Li, R. J., Collins, D. R., Ferrare, R. A., and Brackett, V. G.: Application of aerosol  
407 hygroscopicity measured at the Atmospheric Radiation Measurement Program's Southern Great Plains  
408 site to examine composition and evolution, J. Geophys. Res.-Atmos., 111, D05S12,  
409 doi:10.1029/2004JD005448, 10.1029/2004jd005448, 2006.
- 410 Guan, W. J., Zheng, X. Y., Chung, K. F., and Zhong, N. S.: Impact of air pollution on the burden of chronic  
411 respiratory diseases in China: time for urgent action, Lancet, 388, 1939-1951, 2016.
- 412 Guo, S., Hu, M., Guo, Q., Zhang, X., Schauer, J. J., and Zhang, R.: Quantitative evaluation of emission  
413 controls on primary and secondary organic aerosol sources during Beijing 2008 Olympics, Atmospheric



- 414 Chemistry and Physics, 13, 8303-8314, 10.5194/acp-13-8303-2013, 2013.
- 415 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina,  
416 M. J., and Zhang, R.: Elucidating severe urban haze formation in China, Proceedings of the National  
417 Academy of Sciences of the United States of America, 111, 17373-17378, 10.1073/pnas.1419604111,  
418 2014.
- 419 Gysel, M., Crosier, J., Topping, D. O., Whitehead, J. D., Bower, K. N., Cubison, M. J., Williams, P. I.,  
420 Flynn, M. J., McFiggans, G. B., and Coe, H.: Closure study between chemical composition and  
421 hygroscopic growth of aerosol particles during TORCH2, Atmospheric Chemistry and Physics, 7, 6131-  
422 6144, 2007.
- 423 Gysel, M., McFiggans, G. B., and Coe, H.: Inversion of tandem differential mobility analyser (TDMA)  
424 measurements, Journal of Aerosol Science, 40, 134-151, 10.1016/j.jaerosci.2008.07.013, 2009.
- 425 Han, B., Zhang, R., Yang, W., Bai, Z., Ma, Z., and Zhang, W.: Heavy haze episodes in Beijing during  
426 January 2013: Inorganic ion chemistry and source analysis using highly time-resolved measurements  
427 from an urban site, Science of The Total Environment, 544, 319-329,  
428 <http://dx.doi.org/10.1016/j.scitotenv.2015.10.053>, 2016.
- 429 Heal, M. R., Kumar, P., and Harrison, R. M.: Particles, air quality, policy and health, Chemical Society  
430 Reviews, 41, 6606-6630, 10.1039/c2cs35076a, 2012.
- 431 Hu, M., Peng, J., Sun, K., Yue, D., Guo, S., Wiedensohler, A., and Wu, Z.: Estimation of size-resolved





432 ambient particle density based on the measurement of aerosol number, mass, and chemical size  
433 distributions in the winter in Beijing, *Environ Sci Technol*, 46, 9941-9947, 10.1021/es204073t, 2012.

434 Hu, Q. Q., Fu, H. B., Wang, Z. Z., Kong, L. D., Chen, M. D., and Chen, J. M.: The variation of  
435 characteristics of individual particles during the haze evolution in the urban Shanghai atmosphere,  
436 *Atmospheric Research*, 181, 95-105, 10.1016/j.atmosres.2016.06.016, 2016.

437 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J. G.,  
438 Platt, S. M., Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G.,  
439 Piazzalunga, A., Schwikowski, M., Abbaszade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat,  
440 S., Baltensperger, U., El Haddad, I., and Prevot, A. S. H.: High secondary aerosol contribution to  
441 particulate pollution during haze events in China, *Nature*, 514, 218-222, 10.1038/nature13774, 2014.

442 IPCC: *Climate Change 2013: The Physical Science Basis*, Cambridge, UK, 2013.

443 Khalizov, A. F., Xue, H., Wang, L., Zheng, J., and Zhang, R.: Enhanced Light Absorption and Scattering  
444 by Carbon Soot Aerosol Internally Mixed with Sulfuric Acid, *Journal of Physical Chemistry A*, 113,  
445 1066-1074, 10.1021/jp807531n, 2009.

446 Kulmala, M., Petaja, T., Nieminen, T., Sipila, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto,  
447 P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E., Laaksonen, A., and Kerminen, V. M.:  
448 Measurement of the nucleation of atmospheric aerosol particles, *Nature protocols*, 7, 1651-1667,  
449 10.1038/nprot.2012.091, 2012.



- 450 Levy, M. E., Zhang, R. Y., Khalizov, A. F., Zheng, J., Collins, D. R., Glen, C. R., Wang, Y., Yu, X. Y.,  
451 Luke, W., Jayne, J. T., and Olaguer, E.: Measurements of submicron aerosols in Houston, Texas during  
452 the 2009 SHARP field campaign, *J. Geophys. Res.-Atmos.*, 118, 10518-10534, 10.1002/jgrd.50785, 2013.
- 453 Levy, M. E., Zhang, R. Y., Zheng, J., Tan, H. B., Wang, Y., Molina, L. T., Takahama, S., Russell, L. M.,  
454 and Li, G. H.: Measurements of submicron aerosols at the California-Mexico border during the Cal-Mex  
455 2010 field campaign, *Atmospheric Environment*, 88, 308-319, 10.1016/j.atmosenv.2013.08.062, 2014.
- 456 Li, J. J., Wang, G. H., Ren, Y. Q., Wang, J. Y., Wu, C., Han, Y. N., Zhang, L., Cheng, C. L., and Meng, J.  
457 J.: Identification of chemical compositions and sources of atmospheric aerosols in Xi'an, inland China  
458 during two types of haze events, *Science of the Total Environment*, 566, 230-237,  
459 10.1016/j.scitotenv.2016.05.057, 2016.
- 460 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H., Bi, X. H., Cheng,  
461 P., and Zhou, Z.: Real time bipolar time-of-flight mass spectrometer for analyzing single aerosol particles,  
462 *International Journal of Mass Spectrometry*, 303, 118-124, 10.1016/j.ijms.2011.01.017, 2011.
- 463 Li, P., Yan, R., Yu, S., Wang, S., Liu, W., and Bao, H.: Reinstate regional transport of PM<sub>2.5</sub> as a major  
464 cause of severe haze in Beijing, *Proceedings of the National Academy of Sciences of the United States of*  
465 *America*, 112, E2739-E2740, 10.1073/pnas.1502596112, 2015.
- 466 Lin, Y., Huang, K., Zhuang, G., Fu, J. S., Wang, Q., Liu, T., Deng, C., and Fu, Q.: A multi-year evolution  
467 of aerosol chemistry impacting visibility and haze formation over an Eastern Asia megacity, Shanghai,



- 468 Atmospheric Environment, 92, 76-86, 10.1016/j.atmosenv.2014.04.007, 2014.
- 469 Massling, A., Stock, M., Wehner, B., Wu, Z. J., Hu, M., Brüggemann, E., Gnauk, T., Herrmann, H., and  
470 Wiedensohler, A.: Size segregated water uptake of the urban submicrometer aerosol in Beijing,  
471 Atmospheric Environment, 43, 1578-1589, 2009.
- 472 Momenimovahed, A., and Olfert, J. S.: Effective density and volatility of particles emitted from gasoline  
473 direct injection vehicles and implications for particle mass measurement, Aerosol Sci. Technol., 49, 1051-  
474 1062, 10.1080/02786826.2015.1094181, 2015.
- 475 Niu, H. Y., Hu, W., Zhang, D. Z., Wu, Z. J., Guo, S., Pian, W., Cheng, W. J., and Hu, M.: Variations of  
476 fine particle physiochemical properties during a heavy haze episode in the winter of Beijing, Science of  
477 the Total Environment, 571, 103-109, 10.1016/j.scitotenv.2016.07.147, 2016.
- 478 Olfert, J. S., Symonds, J. P. R., and Collings, N.: The effective density and fractal dimension of particles  
479 emitted from a light-duty diesel vehicle with a diesel oxidation catalyst, Journal of Aerosol Science, 38,  
480 69-82, 10.1016/j.jaerosci.2006.10.002, 2007.
- 481 Pagels, J., Khalizov, A. F., McMurry, P. H., and Zhang, R. Y.: Processing of Soot by Controlled Sulphuric  
482 Acid and Water Condensation Mass and Mobility Relationship, Aerosol Sci. Technol., 43, 629-640,  
483 10.1080/02786820902810685, 2009.
- 484 Park, K., F. C., Kittelson, D. B., and McMurry, P. H.: Relationship between particle mass and mobility  
485 for diesel exhaust particles Environmental Science and Technology, 37, 577-583, 2003.



- 486 Peng, J., Hu, M., Guo, S., Du, Z., Zheng, J., Shang, D., Zamora, M. L., Zeng, L., Shao, M., Wu, Y.-S.,  
487 Zheng, J., Wang, Y., Glen, C. R., Collins, D. R., Molina, M. J., and Zhang, R.: Markedly enhanced  
488 absorption and direct radiative forcing of black carbon under polluted urban environments, Proceedings  
489 of the National Academy of Sciences of the United States of America, 113, 4266-4271,  
490 10.1073/pnas.1602310113, 2016.
- 491 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and  
492 cloud condensation nucleus activity, Atmospheric Chemistry and Physics, 7, 1961-1971, 2007.
- 493 Qiao, T., Zhao, M., Xiu, G., and Yu, J.: Simultaneous monitoring and compositions analysis of PM<sub>1</sub> and  
494 PM<sub>2.5</sub> in Shanghai: Implications for characterization of haze pollution and source apportionment, The  
495 Science of the total environment, 557-558, 386-394, 10.1016/j.scitotenv.2016.03.095, 2016.
- 496 Shi, Y., Chen, J., Hu, D., Wang, L., Yang, X., and Wang, X.: Airborne submicron particulate (PM<sub>1</sub>)  
497 pollution in Shanghai, China: Chemical variability, formation/dissociation of associated semi-volatile  
498 components and the impacts on visibility, Science of the Total Environment, 473, 199-206,  
499 10.1016/j.scitotenv.2013.12.024, 2014.
- 500 Spencer, M. T., Shields, L. G., and Prather, K. A.: Simultaneous measurement of the effective density and  
501 chemical composition of ambient aerosol particles, Environmental Science & Technology, 41, 1303-1309,  
502 10.1021/es061425+, 2007.
- 503 Sun, Y. L., Chen, C., Zhang, Y. J., Xu, W. Q., Zhou, L. B., Cheng, X. L., Zheng, H. T., Ji, D. S., Li, J.,



- 504 Tang, X., Fu, P. Q., and Wang, Z. F.: Rapid formation and evolution of an extreme haze episode in  
505 Northern China during winter 2015, *Scientific Reports*, 6, 10.1038/srep27151, 2016.
- 506 Swietlicki, E., Hansson, H. C., Hameri, K., Svenningsson, B., Massling, A., McFiggans, G., McMurry, P.  
507 H., Petaja, T., Tunved, P., Gysel, M., Topping, D., Weingartner, E., Baltensperger, U., Rissler, J.,  
508 Wiedensohler, A., and Kulmala, M.: Hygroscopic properties of submicrometer atmospheric aerosol  
509 particles measured with H-TDMA instruments in various environments - a review, *Tellus Ser. B-Chem.*  
510 *Phys. Meteorol.*, 60, 432-469, 10.1111/j.1600-0889.2008.00350.x, 2008.
- 511 Tao, Y., Ye, X. N., Ma, Z., Xie, Y. Y., Wang, R. Y., Chen, J. M., Yang, X., and Jiang, S. Q.: Insights into  
512 different nitrate formation mechanisms from seasonal variations of secondary inorganic aerosols in  
513 Shanghai, *Atmospheric Environment*, 145, 1-9, 10.1016/j.atmosenv.2016.09.012, 2016.
- 514 Topping, D. O., McFiggans, G. B., and Coe, H.: A curved multi-component aerosol hygroscopicity model  
515 framework: Part 1 - Inorganic compounds, *Atmospheric Chemistry and Physics*, 5, 1205-1222, 2005.
- 516 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Zamora, M. L., Hu, M., Lin, Y., Peng, J., Guo, S., Meng,  
517 J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian,  
518 P., Marrero-Ortiz, W., Secret, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y.,  
519 Wang, Y., Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D., Liss, P. S.,  
520 Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation from London Fog to Chinese  
521 haze, *Proceedings of the National Academy of Sciences of the United States of America*, 113, 13630-



- 522 13635, 10.1073/pnas.1616540113, 2016.
- 523 Wang, H., Xu, J., Zhang, M., Yang, Y., Shen, X., Wang, Y., Chen, D., and Guo, J.: A study of the  
524 meteorological causes of a prolonged and severe haze episode in January 2013 over central-eastern China,  
525 Atmospheric Environment, 98, 146-157, 10.1016/j.atmosenv.2014.08.053, 2014a.
- 526 Wang, Y., Wan, Q., Meng, W., Liao, F., Tan, H., and Zhang, R.: Long-term impacts of aerosols on  
527 precipitation and lightning over the Pearl River Delta megacity area in China, Atmospheric Chemistry  
528 and Physics, 11, 12421-12436, 10.5194/acp-11-12421-2011, 2011.
- 529 Wang, Y., Khalizov, A., Levy, M., and Zhang, R.: New Directions: Light absorbing aerosols and their  
530 atmospheric impacts, Atmospheric Environment, 81, 713-715, 10.1016/j.atmosenv.2013.09.034, 2013.
- 531 Wang, Y., Zhang, Q., Jiang, J., Zhou, W., Wang, B., He, K., Duan, F., Zhang, Q., Philip, S., and Xie, Y.:  
532 Enhanced sulfate formation during China's severe winter haze episode in January 2013 missing from  
533 current models, J. Geophys. Res.-Atmos., 119, 10.1002/2013jd021426, 2014b.
- 534 Wang, Y., Zhang, R., and Saravanan, R.: Asian pollution climatically modulates mid-latitude cyclones  
535 following hierarchical modelling and observational analysis, Nature Communications, 5,  
536 10.1038/ncomms4098, 2014c.
- 537 Wang, Y. H., Liu, Z. R., Zhang, J. K., Hu, B., Ji, D. S., Yu, Y. C., and Wang, Y. S.: Aerosol  
538 physicochemical properties and implications for visibility during an intense haze episode during winter  
539 in Beijing, Atmospheric Chemistry and Physics, 15, 3205-3215, 10.5194/acp-15-3205-2015, 2015.



- 540 Wu, G., Li, Z., Fu, C., Zhang, X., Zhang, R., Zhang, R., Zhou, T., Li, J., Li, J., Zhou, D., Wu, L., Zhou,  
541 L., He, B., and Huang, R.: Advances in studying interactions between aerosols and monsoon in China,  
542 Science China-Earth Sciences, 59, 1-16, 10.1007/s11430-015-5198-z, 2016a.
- 543 Wu, S., Ni, Y., Li, H., Pan, L., Yang, D., Baccarelli, A. A., Deng, F., Chen, Y., Shima, M., and Guo, X.:  
544 Short-term exposure to high ambient air pollution increases airway inflammation and respiratory  
545 symptoms in chronic obstructive pulmonary disease patients in Beijing, China, Environment International,  
546 94, 76-82, 10.1016/j.envint.2016.05.004, 2016b.
- 547 Xiao, S., Wang, Q. Y., Cao, J. J., Huang, R. J., Chen, W. D., Han, Y. M., Xu, H. M., Liu, S. X., Zhou, Y.  
548 Q., Wang, P., Zhang, J. Q., and Zhan, C. L.: Long-term trends in visibility and impacts of aerosol  
549 composition on visibility impairment in Baoji, China, Atmospheric Research, 149, 88-95,  
550 10.1016/j.atmosres.2014.06.006, 2014.
- 551 Xiao, S., Wang, M. Y., Yao, L., Kulmala, M., Zhou, B., Yang, X., Chen, J. M., Wang, D. F., Fu, Q. Y.,  
552 Worsnop, D. R., and Wang, L.: Strong atmospheric new particle formation in winter in urban Shanghai,  
553 China, Atmospheric Chemistry and Physics, 15, 1769-1781, 10.5194/acp-15-1769-2015, 2015.
- 554 Xie, Y., Ding, A., Nie, W., Mao, H., Qi, X., Huang, X., Xu, Z., Kerminen, V.-M., Petaja, T., Chi, X.,  
555 Virkkula, A., Boy, M., Xue, L., Guo, J., Sun, J., Yang, X., Kulmala, M., and Fu, C.: Enhanced sulfate  
556 formation by nitrogen dioxide: Implications from in situ observations at the SORPES station, J. Geophys.  
557 Res.-Atmos., 120, 12679-12694, 10.1002/2015jd023607, 2015.



- 558 Yang, L., Zhou, X., Wang, Z., Zhou, Y., Cheng, S., Xu, P., Gao, X., Nie, W., Wang, X., and Wang, W.:  
559 Airborne fine particulate pollution in Jinan, China: Concentrations, chemical compositions and influence  
560 on visibility impairment, *Atmospheric Environment*, 55, 506-514, 10.1016/j.atmosenv.2012.02.029, 2012.
- 561 Ye, X. N., Chen, T. Y., Hu, D. W., Yang, X., Chen, J. M., Zhang, R. Y., Khakuziv, A. F., and Wang, L.: A  
562 multifunctional HTDMA system with a robust temperature control, *Advances in Atmospheric Sciences*,  
563 26, 1235-1240, 10.1007/s00376-009-8134-3, 2009.
- 564 Ye, X. N., Ma, Z., Hu, D. W., Yang, X., and Chen, J. M.: Size-resolved hygroscopicity of submicrometer  
565 urban aerosols in Shanghai during wintertime, *Atmospheric Research*, 99, 353-364, 2010.
- 566 Ye, X. N., Ma, Z., Zhang, J. C., Du, H. H., Chen, J. M., Chen, H., Yang, X., Gao, W., and Geng, F. H.:  
567 Important role of ammonia on haze formation in Shanghai, *Environ Res Lett*, 6, Artn 024019  
568 Doi 10.1088/1748-9326/6/2/024019, 2011.
- 569 Ye, X. N., Tang, C., Yin, Z., Chen, J. M., Ma, Z., Kong, L. D., Yang, X., Gao, W., and Geng, F. H.:  
570 Hygroscopic growth of urban aerosol particles during the 2009 Mirage-Shanghai Campaign, *Atmospheric  
571 Environment*, 64, 263-269, 10.1016/j.atmosenv.2012.09.064, 2013.
- 572 Yin, Z., Ye, X., Jiang, S., Tao, Y., Shi, Y., Yang, X., and Chen, J.: Size-resolved effective density of urban  
573 aerosols in Shanghai, *Atmospheric Environment*, 100, 133-140, 10.1016/j.atmosenv.2014.10.055, 2015.
- 574 Zhang, Q., Quan, J., Tie, X., Li, X., Liu, Q., Gao, Y., and Zhao, D.: Effects of meteorology and secondary  
575 particle formation on visibility during heavy haze events in Beijing, China, *Science of The Total*





- 576 Environment, 502, 578-584, <http://dx.doi.org/10.1016/j.scitotenv.2014.09.079>, 2015a.
- 577 Zhang, R., Khalizov, A., Wang, L., Hu, M., and Xu, W.: Nucleation and Growth of Nanoparticles in the  
578 Atmosphere, Chemical Reviews, 112, 1957-2011, 10.1021/cr2001756, 2012.
- 579 Zhang, R., Guo, S., Zamora, M. L., and Hu, M.: Reply to Li et al.: Insufficient evidence for the  
580 contribution of regional transport to severe haze formation in Beijing, Proceedings of the National  
581 Academy of Sciences of the United States of America, 112, E2741-E2741, 10.1073/pnas.1503855112,  
582 2015b.
- 583 Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.:  
584 Formation of Urban Fine Particulate Matter, Chemical Reviews, 115, 3803-3855,  
585 10.1021/acs.chemrev.5b00067, 2015c.
- 586 Zhang, R. Y., Suh, I., Zhao, J., Zhang, D., Fortner, E. C., Tie, X. X., Molina, L. T., and Molina, M. J.:  
587 Atmospheric new particle formation enhanced by organic acids, Science, 304, 1487-1490,  
588 10.1126/science.1095139, 2004.
- 589 Zhang, R. Y., Khalizov, A. F., Pagels, J., Zhang, D., Xue, H. X., and McMurry, P. H.: Variability in  
590 morphology, hygroscopicity, and optical properties of soot aerosols during atmospheric processing,  
591 Proceedings of the National Academy of Sciences of the United States of America, 105, 10291-10296,  
592 2008.
- 593 Zhao, J., Khalizov, A., Zhang, R., and McGraw, R.: Hydrogen-Bonding Interaction in Molecular



- 594 Complexes and Clusters of Aerosol Nucleation Precursors, *Journal of Physical Chemistry A*, 113, 680-  
595 689, 10.1021/jp806693r, 2009.
- 596 Zhao, X. J., Zhao, P. S., Xu, J., Meng, W., Pu, W. W., Dong, F., He, D., and Shi, Q. F.: Analysis of a winter  
597 regional haze event and its formation mechanism in the North China Plain, *Atmospheric Chemistry and*  
598 *Physics*, 13, 5685-5696, 10.5194/acp-13-5685-2013, 2013.
- 599 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T.,  
600 Chang, D., Poeschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the  
601 impact of synoptic weather, regional transport and heterogeneous reactions, *Atmospheric Chemistry and*  
602 *Physics*, 15, 2969-2983, 10.5194/acp-15-2969-2015, 2015.
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604 **Figure and Table Captions**

605 Figure 1 Schematic diagram of HTDMA-APM system.

606 Figure 2. Temporal evolutions of  $PM_{1.0}$ ,  $PM_{2.5}$ , and  $PM_{10}$  concentrations during the winter observation.

607 Figure 3 Variations of sulfate, nitrate, and ammonium concentrations as a function of  $PM_{1.0}$  mass loading.

608 Figure 4 Box plots showing hygroscopicity parameter and effective density at each dry diameter over the  
609 whole observation. The whiskers represent the 5<sup>th</sup> and 95<sup>th</sup> percentile, the two borders of box display the  
610 25<sup>th</sup> and 75<sup>th</sup> percentile, and the band in each box denotes the median.

611 Figure 5 Temporal evolutions of particle number size distribution (A), volume size distribution (B), total  
612 number concentration and total volume concentration (C), and  $PM_{1.0}$  concentration and calculated PM  
613 (less than 600 nm in mobility diameter) concentration during the representative PM episode from 7 to 12  
614 January.

615 Figure 6 Evolutions of particle hygroscopic growth factor and effective density for different sizes during  
616 the representative PM episode.

617 Figure 7 Temporal evolutions of chemical compositions determined by SPAMS during the representative  
618 PM episode.

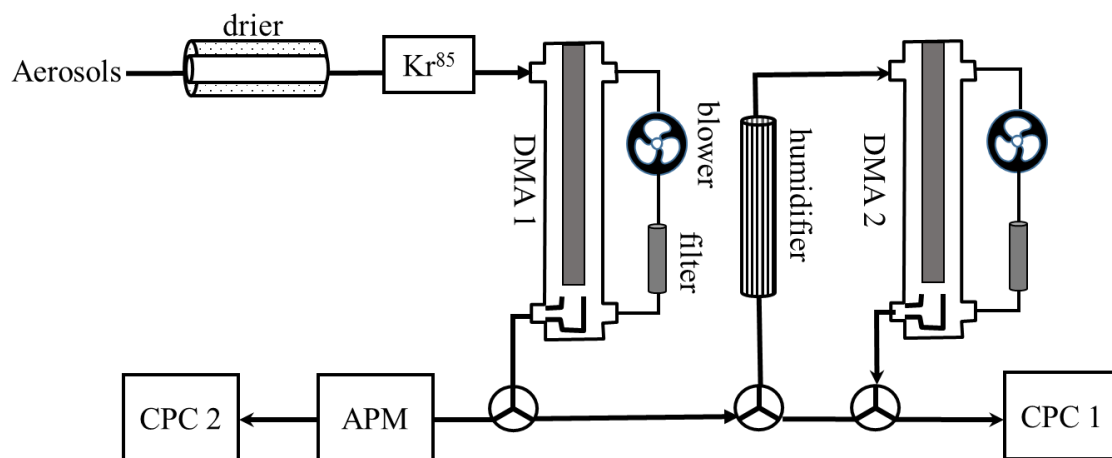
619 Figure 8 Particle hygroscopicity and density during the two particle growth processes

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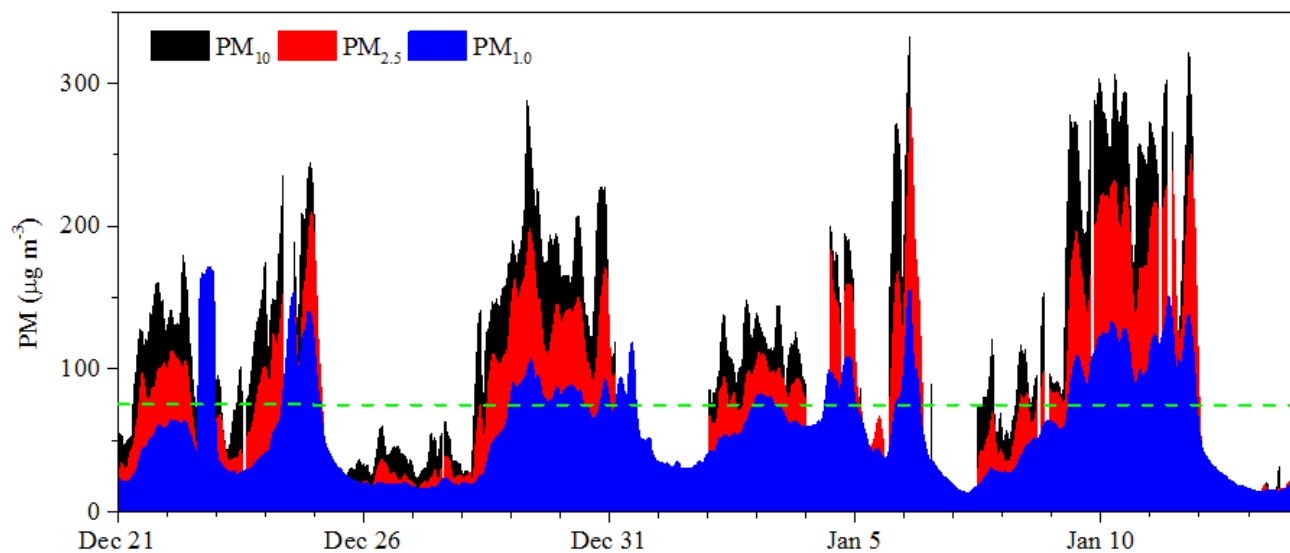


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Figure 1 Schematic diagram of HTDMA-APM system.



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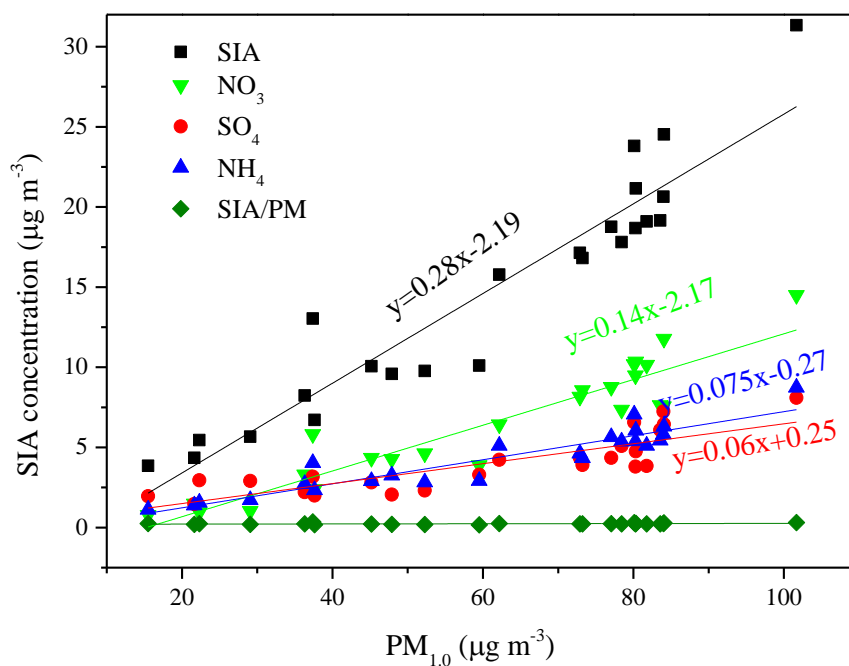
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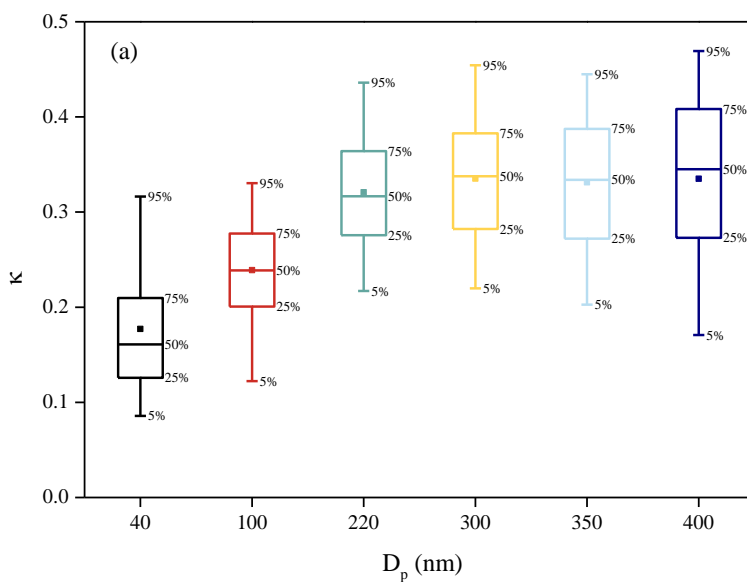
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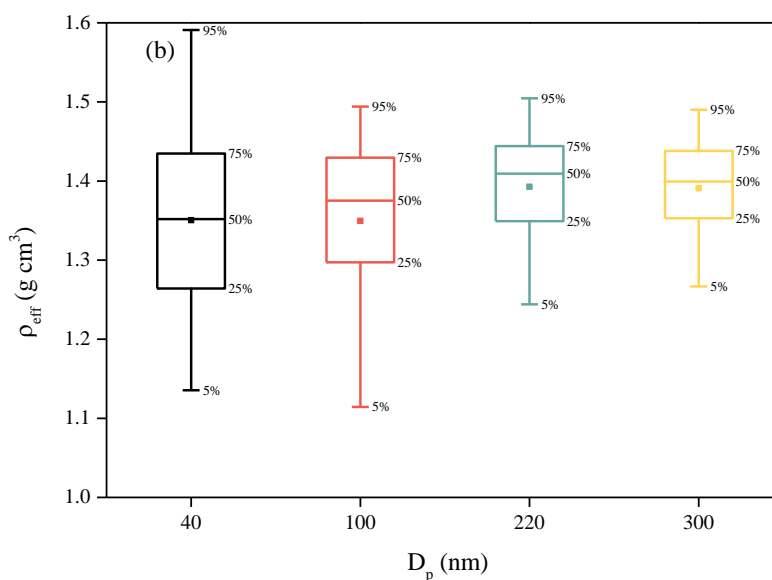
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636 Figure 3 Variations of sulfate, nitrate, and ammonium concentrations as a function of  $PM_{1.0}$  mass loading

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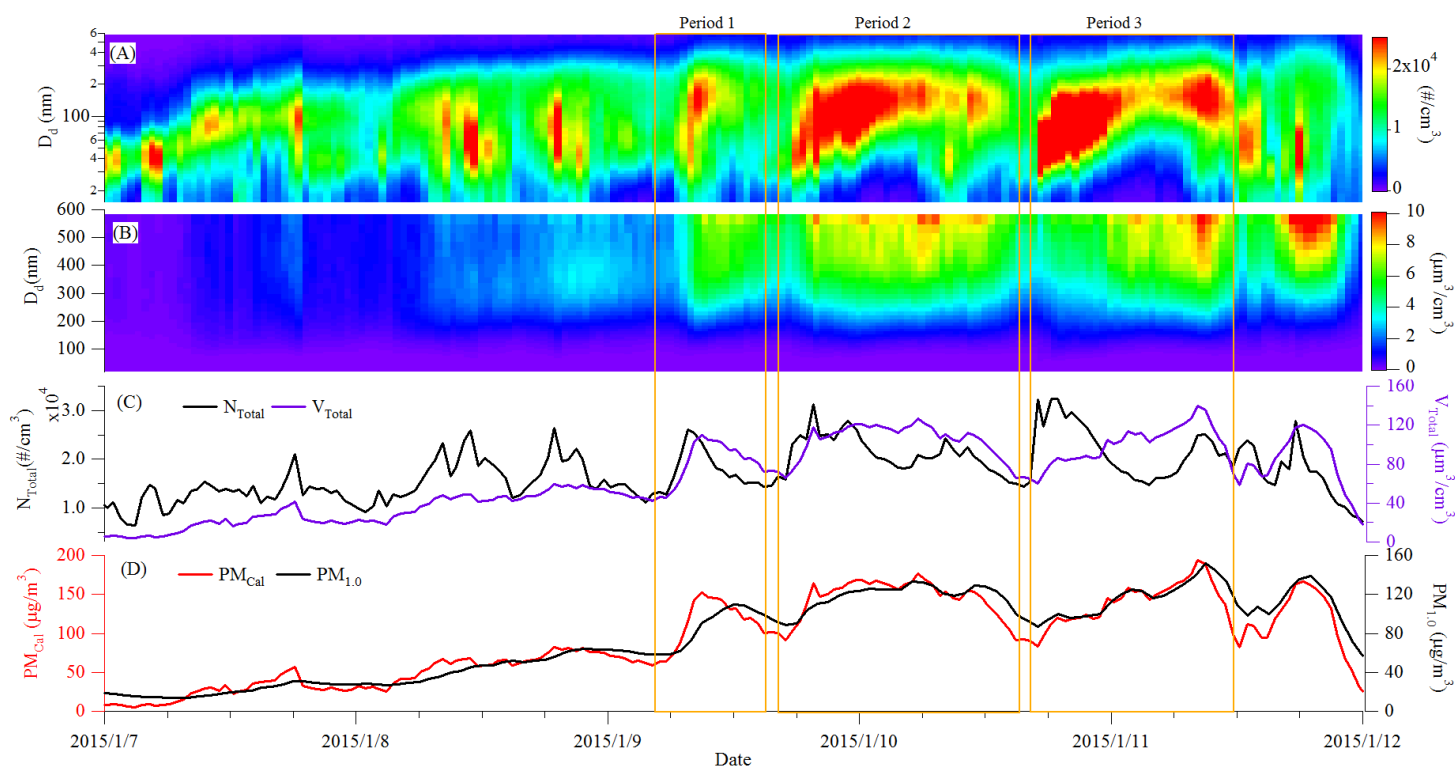
640 Figure 4 Box plots showing hygroscopicity parameter and effective density at each dry diameter over the  
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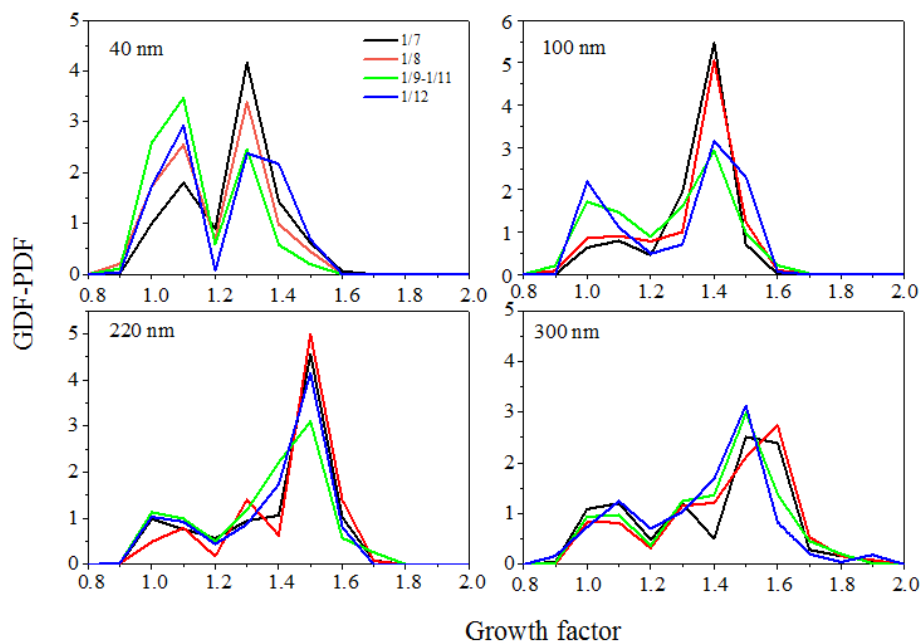
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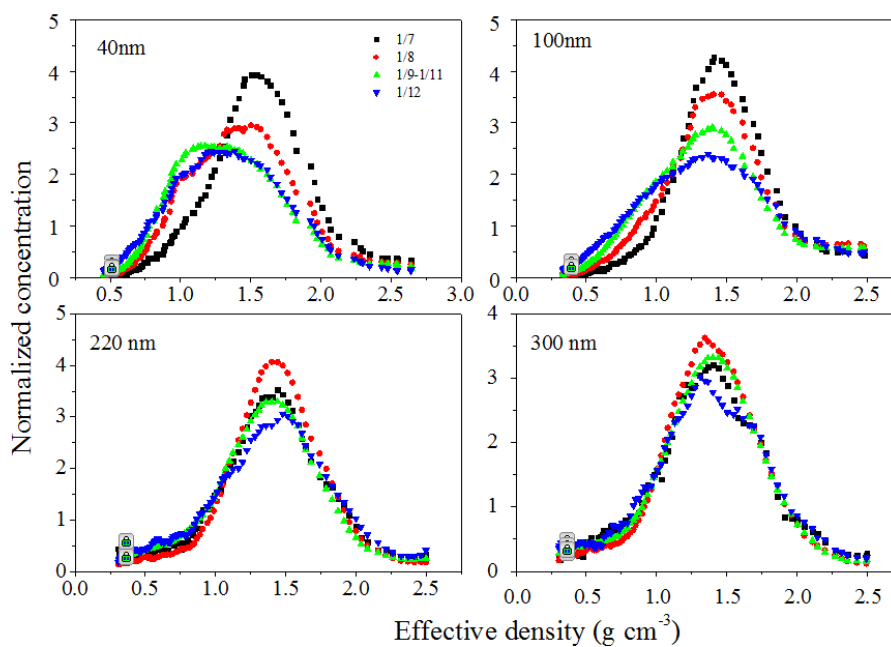
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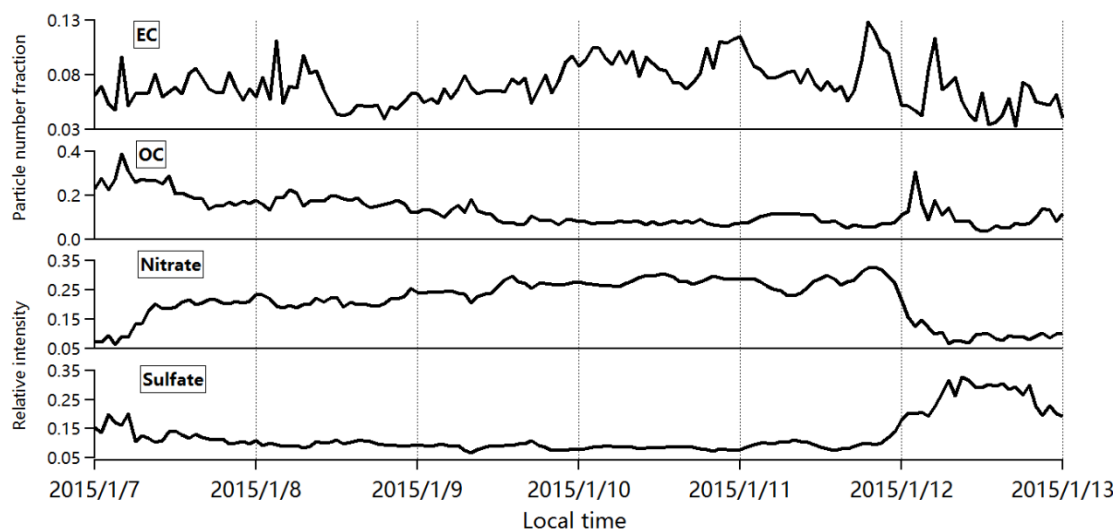


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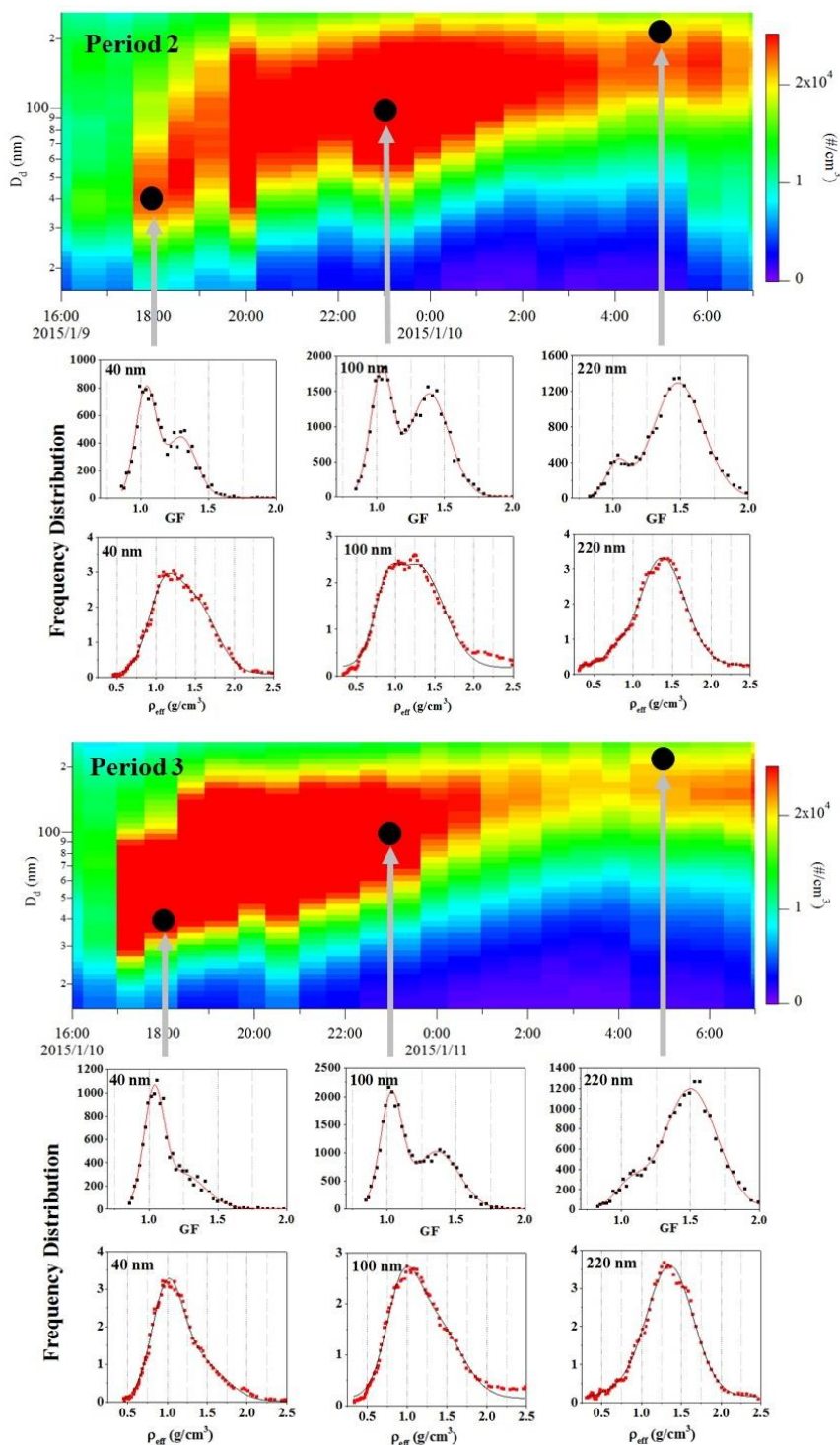
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