## **Final Author comments**

Authors' response to Referee #1, Referee #2 and Editor's comments on "UV measurements at Marambio and Ushuaia during 2000-2010" by Kaisa Lakkala et al.

The authors thank the Referees and the Editor for constructive comments and reply to all comments here below. The answer is structured as follow: (1) comments from Referee/Editor, (2) author's response, (3) author's changes in the manuscript.

## Referee #1

(1) The authors should give an estimate of the absolute uncertainty of the presented UV index data.
 (2) To estimate the uncertainty of the presented UV index data, we calculated the combined uncertainty and used the coverage factor of 2.

(3) The text has been updated to be at Page 10, lines 15-19: "Using the a step of 5% as uncertainty related to the drift of the calibration of the site instrument between calibrations, the drift of the traveling reference between calibrations (3% for 2004–2006), the uncertainty of the calibration of the traveling reference (Lakkala et al., 2005), the uncertainties related to the transfer of the irradiance scale (Lakkala et al., 2005) and the uncertainty of the method to calculated the dose rates (Dahlback, 1996), the combined uncertainty was calculated to be 9.5%. The expanded uncertainty was 19% using a coverage factor of 2 (Table 3).

(1) The authors state correctly that for clear sky conditions the solar zenith angle, the ozone amount and the coverage of the ground with snow are the most important parameters. To show the effect of ozone and snow more specifically, an additional figure with the time series of the UV index at a fixed solar zenith angle (e.g.  $70^{\circ}$  or  $75^{\circ}$ ) could be impressive. Then, potentially, the relation between the UV index and ozone could be interpreted not only qualitatively, but also quantitatively (e.g. in terms of the radiation amplification factor).

(2) The UV indices at SZA between 69.95 and 70.05 were selected to study the impact of snow in Marambio. The impact of snow was not supposed to be clearly seen, as there is snow on the ground almost year round. The amount of snow may vary depending on the weather conditions, but the highest snow amount occurs during winter and the smallest amount during summer months. Clear days were selected. For the whole time series, the RAF was calculated using the power law and it was found to be 1.16. For biggest ozone deviations, this RAF underestimated the effect in UV radiation, where RAF of around 1.3 were seen. Biggest ozone variations occurred in October, when there most likely snow in the ground at the station of Marambio.

(3) The plot of the RAF was including in the manuscript together with the corresponding text: P. 4: "The radiation amplification factor (RAF) was calculated for erythemally weighted UV irradiances measured at SZA between 69.95 and 70.05 in Marambio. Clear sky data was selected using the cloud transmission factor (CLT) calculated from the NILU- UV irradiance at 380 nm (Høiskar et al., 2003). The power law (e.g., Booth and Madronich (1994)) was used to calculate the RAF (Eq. (5)). RAF =ln(O 3 /O 3 \* )/ln(U V \* /U V ) (5)

where O 3 \* is the first total ozone value and O 3 the second, respectively. U V \* is the first erythemally weighted UV irradiance and U V the second, respectively.

P 13: "The RAF was calculated to be 1.16, when excluding the winter months May-July from the analysis. When largest ozone decreases occured, the RAF could exceed 1.25 (Fig. 9) suggesting an additional increase due to snow on the ground."

(1) I wonder that in Fig. 7a and 10a the variability of the ratios maximum/mean and minimum/mean is quite similar throughout the year, even on the days around winter solstice, when the absolute level of the mean is extremely small. Therefore very small differences at this time should give much higher variability of the ratio compared to the other days of the year.

(2) The calculations were checked and no mistakes were found. We agree that the climatological mean of daily maximum UV index and daily dose is very small around the winter solstice, but also the variability is small during that period, as the sun is near the horizon (high SZA) and low intensity. The variability of the rations maximum/mean and minimum/mean ranged from 1.2 to 1.8 and from 0.001 to 0.85, respectively. During the other days of the year, the variability can be as high, e.g. for rainy days, when the minimum is very low.

(3) No update was made to the text.

Technical details:

(1) Several times the term "daily mean" is used, but sometimes this is a bit misleading (e.g. p.4, l.8 or legend of Fig. 8). "Daily mean" should not be understood as the mean over the day, but as the day-to-day variation of the mean, where the mean is calculated from the respective days of each year of observations.

(2) The authors have checked the text regarding the use of the term "daily mean". And changed the wording to clarify the text.

(3) The following sentences have been updated: P4 l. 14-18"For each day of the year, means of daily maximum UV index and daily dose were calculated from the respective days of each year of observation. From now on in this paper, these means are called "daily climatology", even if they are not real climatological means from statistical point of view as the time period is too short. Daily deviations (DEV) as percentage were calculated from this daily climatology following Eq. (4), where X is the daily maximum UV index or daily dose and "clim" is the daily climatology."

(2) From that paragraph onwards, the "daily mean" which was calculated from the respective days of each year of observations was called "daily climatology".

(1) p.13, l. 5: "did not reach" instead of "did not reached". (2,3) Changed.

(1) p. 17, l. 8: "Fig. 9" instead of "Figs. 9 and 10".

(2) Changed.

(3) The reference to the Fig. 10 (now Fig. 11) is added to the following sentence: "The high fluctuation in maximum values (Fig. 11) is due to the impact of year to year changes in cloudiness and total ozone."

## Referee #2

(1) p1, l. 7: "daily doses" instead of "dailydoses", pls change throughout the document

(2,3) Changed throughout the document.

(1) p2,l.13: "begining" -> "beginning"(2,3) Changed.

(1) Capitalize references to specific figures, tables, chapters, sections, equations: p3, l.29: "eq. 1" -> "Eq. 1". Not sure about the specific style guides of ACP, but I suggest to capitalize e.g. Equation 1, Figure 1, Appendix C since they are names of the entities they refer to. Please unify throughout the manuscript.

(2,3) References to specific figures, tables, chapters, sections and equations have been capitalized and changed to follow the style guides of ACP.

(1) p4, l20ff: you present lamp measurements in great detail just to inform the reader four sites later on p9, l6. that they are not used for the instrument calibration. I recommend giving that information already in Section 2.2.1.

(2) The information has been moved earlier.

(3) The text has been updated and is now: Page 5, line 3 " The results of the lamp tests were used to confirm the studied changes in the scaling coefficients, but not used in the absolute calibration."

(1) p9, l2: You state that the instruments are calibrated 2-3 times a year, some channels show a significant drift during that time scale. How big are the steps that occur in your data after those calibrations? Provide a percentage (up to x

(2) The percentage has been provided and added to Page 10 line 14.

(3) The following text was added to the manuscripts: "The highest step between calibrations occured in Marambio in January 2000 and it was 11% for erythemally weighted UV dose rates."

(1) p9, l8-10 channel 2 -> channel two (2,3) Changed.

(1) Results: Your results focus on the presentation of long-term time series and daily means - did you consider mentioning the number of measurements performed per day (or the time intervals between single measurements)? You could for example show data of one particular day to illustrate the capabilities of tracking fast ozone changes. If the paper gets too lengthy consider skipping some figures of UV daily dose since they show a very similar behavior like those of the UV index.

(2) The following sentence already describes the frequency of recored measurements per day: P3 l 11 " In routine operation, one minute averages are recorded by a built-in data logger." We added a plot (Figure 6) of the record high UV index day, 18.11.2007 in Marambio to illustrate the capabilities of tracking fast ozone/UV changes. And added the following sentence on P3 "This allows the detection of rapid changes in cloudiness and total ozone amounts."

(3) The text is now:

P3: "In routine operation, one minute averages are recorded by a built-in data logger. This allows the detection of rapid changes in cloudiness and total ozone amounts." and

P14, l. 5: At Marambio, extreme (i.e., UV indices of 11 or more, WMO (1997)) daily maximum UV indices were measured on 5 days, and the peak value of 12 was measured on 18 Nov 2007. The daily variation of the UV index on that record day is shown in Fig. 6. The total ozone amount on that day was 213 DU.

## Editor

Figures 1 and 2: Please revise the layout of panels. You could a layout similar to that of Fig. 3 for the six channels and a fourth row showing the temperature and legend. (2,3) The Figures have been revised and updated as suggested by the Editor.

Figure 4: Change caption to: Daily maximum UV index. Vertical axes labels are too dense. Either increase the axes height, or reduce the number of labels. (2,3) The caption has been changed and the axes height has been increased.

Figure 5. You may consider changing the data to kJ/m2; the vertical axes labels with thousands look quite ugly. Same for Figure 11. (2,3) Changed to  $kJ/m^2$ 

Figures 7 and 8: Labels are too small.

(2,3) Labels have been increased. The final size of the labels will depend on the size of the figure in the final version.

All Figures: Please maintain consistency in the size of labels for all figures. (2,3) The figures have been harmonized regarding the labels.

# UV measurements at Marambio and Ushuaia during 2000-2010

Kaisa Lakkala<sup>1,2</sup>, Alberto Redondas<sup>3</sup>, Outi Meinander<sup>4</sup>, Laura Thölix<sup>4</sup>, Britta Hamari<sup>1</sup>, Antonio Fernando Almansa<sup>3</sup>, Virgilio Carreno<sup>3</sup>, Rosa D. Garcia<sup>5,3,6</sup>, Carlos Torres<sup>3</sup>, Guillermo Deferrari<sup>7,8</sup>, Hector Ochoa<sup>9</sup>, Germar Bernhard<sup>10</sup>, Ricardo Sanchez<sup>11</sup>, and Gerrit de Leeuw<sup>1</sup> <sup>1</sup>Finnish Meteorological Institute, Climate Research Programme, Kuopio, Finland <sup>2</sup>Finnish Meteorological Institute, Space and Earth Observation Centre, Sodankylä, Finland <sup>3</sup>Izaña Atmospheric Research Center, Agencia Estatal de Meteorología, Tenerife, Spain <sup>4</sup>Finnish Meteorological Institute, Climate Research Programme, Helsinki, Finland <sup>5</sup>Air Liquide España, Delegación Canarias, Candelaria, Spain <sup>6</sup>Atmospheric Optics Group, Valladolid University, Valladolid, Spain <sup>7</sup>Centro Austral de Investigaciones Cientificas (CADIC/CONICET), Ushuaia, Argentina <sup>8</sup>Universidad Nacional de Tierra del Fuego, Ushuaia, Argentina <sup>9</sup>Dirección Nacional del Antártico-Instituto Antártico Argentino, Buenos Aires, Argentina <sup>10</sup>Biospherical Instruments, Inc., San Diego, U.S.A <sup>11</sup>Servicio Meteorológico Nacional, Argentina

Correspondence to: Kaisa Lakkala (kaisa.lakkala@fmi.fi)

**Abstract.** Solar ultraviolet (UV) irradiances were measured with NILU-UV multichannel radiometers at Ushuaia ( $54^{\circ}S$ ) and Marambio ( $64^{\circ}S$ ) between 2000 and 2013. The measurements were part of the Antarctic NILU-UV network, which was started in cooperation between Spain, Argentina and Finland. The erythemally weighted UV irradiance time series of both stations

- 5 were analyzed for the first time in this study. The quality assurance procedures included a traveling referce instrument to transfer the irradiance scale to the stations. The time series were homogenized and high quality measurements were available for the period 2000–2010. During this period UV indices of 11 or more were measured on 5 and 35 days at Marambio and Ushuaia, correspondingly. At Marambio, the peak daily maximum UV index of 12 and dailydoses daily doses of around 7  $kJ/m^2$  were measured in November 2007. Typically the The highest UV daily doses were at both stations were typically
- 10 around 6  $kJ/m^2$ , and they occurred on time periods, when the station was and occurred when the stations were inside the polar vortex with, resulting in very low total ozone amount. At both stations, dailydoses daily doses of late November could even exceed those of summer. At Marambio, in some years, also dailydoses daily doses in October can be as high as those during the summer. At Ushuaia, the peak daily maximum UV index of 13 was measured twice: in November 2003 and 2009. Also during those days, the station of Ushuaia was inside the polar vortex.

15 Copyright statement. TEXT

#### 1 Introduction

The NILU-UV Antarctic network was established in 1999/2000 in scientific cooperation between the Spanish Agencia Estatal de Meteorología (AEMET), the Finnish Meteorological Institute (FMI) and the Argentinian Dirección Nacional del Antártico-Instituto Antártico Argentino (DNA-IAA) and Centro Austral de Investigaciones Científicas (CADIC). The goal was to promote

5 research of stratospheric ozone and ultraviolet (UV) radiation in the Antarctic region and to serve multidisciplinary research by providing UV and ozone data.

The well-known Antarctic ozone depletion was first discovered by Farman et al. (1985) in the mid-1980s. Satellite measurements confirmed that the ozone hole existed over whole Antarctica. Soon after, more in situ measurements were started to monitor total ozone and UV radiation. Also regular balloon soundings were launched to monitor the profile of the ozone distri-

- 10 bution in the atmosphere (e.g. Karhu et al. (2003)). The knowledge in the mechanisms of the ozone depletion increased, and the role of ozone depleting subtances and cold temperatures inside the polar vortex was understood (WMO, 1990, 1999; Pazmiño et al., 2005). Therefore the stations of the Antarctic NILU-UV network were chosen regarding their location: Belgrano (77°, 35°W) was mostly located inside, Marambio (64°S, 56°W) at the edge and Ushuaia (54°S, 68°W) outside the polar vortex. NILU-UV multichannel radiometers (Høiskar et al., 2003) were set up at the stations to continuosly monitor total ozone and
- 15 UV irradiance.

The measurements of the network were interrupted in 2014, and the whole UV time series of Ushuaia and Marambio are analysed for the first time in this study. Since the beginning beginning of the network, special attention was given to the quality assurance of the measurements in order to make measurements world wide comparable (Lakkala et al., 2005). A traveling reference was used to transfer the irradiance scale to both Marambio and Ushuaia. As the instrument of Belgrano didn't follow

20 the same quality assurance procedures, albeit it was regularly exchanged with a newly calibrated instrument, the time series of Belgrano is not discussed in this study.

The first measurements of the network started in the middle of the ozone depletion of the Antarctic ozone layer (WMO, 2003, 2007). Later, the recovery of stratospheric ozone was predicted (Andrady et al., 2009), and nowadays, 30 years after signing the Montreal Protocol restricting ozone depleting substances, first signs of the stratospheric ozone recovery have been

25 observed (Solomon et al., 2016). The NILU-UV time series of Ushuaia and Marambio presented in this study are important for the future: The time series serve as reference representing times when severe stratospheric ozone depletion occurred.

Observations at these locations complement measurements with spectroradiometers at the South Pole, McMurdo Station (78°S), Palmer Station (65°S) and Ushuaia, which started in the early 1990s and are part of the U.S. National Science Foundation's (NSF) UV Monitoring Network (Booth et al., 1994; Bernhard et al., 2010). Bernhard et al. (2008) and Bernhard et al.

30 (2010) analyzed the data of these sites and found a large effect of the ozone hole on the UV index at the three Antarctic sites, and to a lesser extent at Ushuaia. Their analysis showed that at South Pole the UV index is on average 20-80% higher during the ozone hole period than between January and March. They studied the UV index compared to estimates of UV index prior the ozone hole period, and found that the average UV index between 1991 and 2006 was 55%–85% higher than the estimates for 1963–1980. The highest UV index measured at Palmer was 14.8. At Ushuaia, UV indices as high as 11.5 were measured

in October, when the ozone hole was over the city. Eleftheratos et al. (2015) studied changes in irradiances at 305 and 325 nm at Southern latitudes  $55^{\circ}$ S-77°S, but didn't find any statistically significant changes between 1990 and 2011.

#### 2 Materials and Methods

20

#### 2.1 NILU-UV radiometers at Ushuaia and Marambio

- 5 The NILU-UV multichannel radiometer measures UV radiation with five channels with center wavelengths at 305, 312, 320, 340 and 380 nm. The full width half maximum (FWHM) of the measured wavelength band is around 10 nm. A sixth channel measures the photosynthetically active radiation (PAR) at wavelengths between 400 and 700 nm. Erythemally weighted irradiances, UVB (280–320 nm) and UVA (320–400 nm) irradiances can be retrieved from the measurements. In addition to UV irradiances, total ozone and information about the cloud optical depth can be calculated.
- 10 The instrument is described in details in Høiskar et al. (2003). It is weatherproof and designed to operate in harsh environment. The instrument has a flat Teflon diffuser, interference filters and silicon detectors. In order to keep the instrument stable at different temperatures, internal heating keep the temperature constant at  $40^{\circ}$ C.

In routine operation, one minute averages are recorded by a built-in data logger. This allows the detection of rapid changes in cloudiness and total ozone amounts. There is a possibility to record measurements each second, and this option was used during

15 lamp measurements. Three weeks of one minute averages can be stored in the memory of the instrument. In the NILU-UV network, the data was however daily/weekly transfered automatically or manually to the servers of AEMET and FMI.

Ushuaia is located in the archipelago of Terra-Tierra del Fuego, Argentina. The NILU-UV instrument was set up on the roof of the CADIC facility next to a SUV-100 spectroradiometer that was part of the U.S. National Science Foundation's Ultraviolet Spectral irradiance Monitoring Network (Booth et al., 1994). The altitude of the station is 23 m. The station is surrounded by mountains, which have snow cover most of the year. In the South, there is the Beagle Channel, which connects Ushuaia to the

South Pacific Ocean. There is snow on the ground from June to September.

The research station of Marambio, Base Marambio II, is located in the point of the North East Antarctic penninsula. At Marambio, the NILU-UV was set up on the roof of the main scientific building. The altitude of the station is 196 m. The climate is characterized by mean temperatures below zero the whole year and strong winds. The temperatures range from

25 -15°C in June to -2° in December (from http://www.smn.gov.ar/serviciosclimaticos/). Fog is frequent, and the most foggy months are December and January. Summer is the cloudiest season and winter have the least clouds. The snow stays on the ground the year round.

#### 2.1.1 Erythemally weighted UV products

The time series of erythemally weighted UV irradiances were calculated using the method described in Dahlback (1996). In 30 According to this method, UV dose rates (D) are calculated from the were derived from measurements of the NILU-UV's five UV channels using a linear function (eq. 1). In the linear function, combination (Eq. (1)) where the raw data (V) measured at  $V_i$ ) measured by each channel (i) was multiplied with a coefficient (a), which was determined during the calibration of the NILU-UV. *i*) are multiplied with coefficients  $(a_i)$ .

$$D = \sum_{i=1}^{5} a_i V_i \tag{1}$$

5 The coefficients  $(a_i)$  of eq. 1 were determined by utilizing a combination of comparisons with a reference spectroradiometer and radiative transfer calculations. The  $a_i$  of Eq. (1) were in turn determined by solving the system of linear equations described by Eq. 7 of Dahlback (1996) (here (Eq. (2)):

$$\sum_{i=1}^{M} a_i k_i \sum_{\lambda=0}^{\infty} R'_{i\lambda} F_{\lambda} = \sum_{\lambda=0}^{\infty} A_{\lambda} F_{\lambda}$$
(2)

The coefficients k<sub>i</sub> are calibration factors. In our implementation, these factors were derived from a vicarious calibration of the
 NILU-UV's raw data with measurements of the spectroradiometer of the Norwegian Radiation Protection Authority (NRPA) was used as reference spectroradiometer (Johnsen et al., 2002). The action spectrum of McKinlay and Diffey (1987) was used, when calculating the coefficients to retrieve erythemally weighted UV dose rates.

$$D = \sum_{i=1}^{5} a_i V_i$$

15

20

(Johnsen et al., 2002), which were weighted with the relative spectral response functions  $R'_{i\lambda}$  of the NILU-UV's channels (Eq. (3) of Dahlback (1996)).  $A_{\lambda}$  is the erythemal action spectrum by McKinlay and Diffey (1987) and  $F_{\lambda}$  are model spectra calculated with the radiative transfer model using pseudo spherical DISORT.

In this work the UV indices and daily doses, calculated from the erythemally weighted dose rates, were studied. The UV index was calculated by multiplying the erythemally weighted dose rate expressed in unit  $W/m^2$  by 40 (e.g., WMO (1997)). The erythemally weighted daily doses (H) were calculated by integrating the erythemally weighted dose rates over the whole day (eq.3Eq. (3), T<sub>1</sub>=00:00 h and T<sub>2</sub>=23:59-24:00 h local time).

$$H = \int_{T_1}^{T_2} D(T) dT \tag{3}$$

For each day of the year, daily-means of daily maximum UV index and daily dose were calulated using the whole time seriescalculated from the respective days of each year of observation. From now on in this paper, these means are called "daily climatology", even if they are not real climatological means from statistical point of view as the time period is too short.

25 Daily deviations (DEV) as percentage were ealulated from these means following eq. 4calculated from this daily climatology following Eq. (4), where X is the daily maximum UV index or daily dose and "mean clim" is the mean value (2000–2010) of that day.

$$\underline{DEV = \frac{(X - mean)}{mean} * 100\%}$$

daily climatology.

$$DEV = \frac{(X - clim)}{clim} * 100\%$$

5 The radiation amplification factor (RAF) was calculated for erythemally weighted UV irradiances measured at SZA between 69.95 and 70.05 in Marambio. Clear sky data was selected using the cloud transmission factor (CLT) calculated from the NILU-UV irradiance at 380 nm (Høiskar et al., 2003). The power law (e.g., Booth and Madronich (1994)) was used to calculate the RAF (Eq. (5)).

$$RAF = \frac{\ln(O_3/O_3^*)}{\ln(UV^*/UV)},\tag{5}$$

10 where  $O_3^*$  is the first total ozone value and  $O_3$  the second, respectively.  $UV^*$  is the first erythemally weighted UV irradiance and UV the second, respectively.

### 2.2 QC of UV data at Marambio and Ushuaia

details in Lakkala et al. (2005) and shortly explained in the next section.

The quality control of the NILU-UVs of Marambio and Ushuaia were performed by the operators at the station. The everyday housekeeping included cleaning of the diffuser, check of leveling and clock adjustment. Regular lamp measurements were performed at the stations and the data was regularly delivered. The quality control of the NILU-UV network is explained in

#### 2.2.1 Lamp measurements

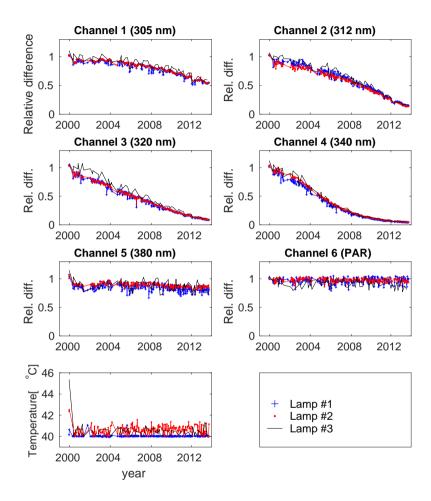
15

Lamp measurements were performed at the station every two weeks. A 100 W lamp was mounted on a lamp holder and placed over the NILU-UV. After the warm up of the lamp, the irradiance was measured during 15 minutes with a time step of one

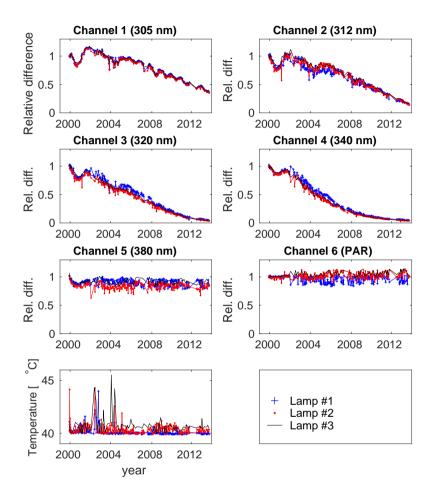
20 second. The temperature of the instrument was also recorded. Each time two lamps were measured and every third time a third lamp was measured in addition. This led to different burning times of the lamps, which made it possible to detect drifts in lamp irradiances.

Each lamp test was compared to the average of the first three measurements at each channels (Fig. 1 and 2). Severe drift in the signal of the instrument was already detected in channels 1–4 after a couple of years of operation of the NILU-UV. The

25 large drifts were caused by the material from which the filters were made: They degraded due to UV radiation. After 4 years of operation, decreases up to 40% in the sensitivity of some filters were observed (Lakkala et al., 2005). The degradation continued through the whole measurement time period. After 2010, the sensitivity of the channels 2–4 decreased more than 80% (Fig. 1 and 2). Due to these large changes, it was no longer possible to correct data after 2010. Until that year, the data was corrected using the scaling of the channels to those of the traveling reference during solar comparisons (see <a href="#">ChapterSection 2.3.2</a>). The results of the lamp tests were used to confirm the studied changes in the scaling coefficients, but not used in the absolute calibration. The data of the years 2011–2013 had to be excluded from the analysis.



**Figure 1.** Lamp measurements at Ushuaia 2000-2013. The raw data of the lamp is divided by the mean of the three first mesurements. The temperature of the NILU-UV during the measurement is also shown.



**Figure 2.** Lamp measurements at Marambio 2000-2013. The raw data of the lamp is divided by the mean of the three first mesurements. The temperature of the NILU-UV during the measurement is also shown.

#### 2.3 QA of the network

5 The quality assurance of the network was based on the use of a traveling reference to transfer the irradiance scale to the NILU-UV radiometers of Ushuaia and Marambio. The details are explained in Lakkala et al. (2005) and the procedures are shortly described in the next sections.

#### 2.3.1 Traveling reference

The traveling reference of the network was a NILU-UV radiometer, which had filters with similar spectral responses as the
NILU-UV radiometers of Ushuaia and Marambio. It was calibrated by the manufacturer in Norway in 1999, 2000, 2001, 2002, 2004, 2005, 2007 and 2010. The irradiance scale was traceable to the National Institute of Standards and Technology (NIST) via the Swedish Testing and Research Institute (SP) (Johnsen et al., 2002).

The quality control of the traveling reference included lamp tests before and after each solar comparison. The quality assurance included comparisons with spectroradiometers in Finland and Ushuaia. In Finland, the traveling reference was compared

- 15 to the Brewer spectroradiometers of FMI at either Sodankylä (67°N) or Jokioinen(61°N). The irradiance scale of the Brewer was traceable via the National Standard Laboratory MIKES, Aalto University, Finland, to SP (Heikkilä et al., 2016; Lakkala et al., 2008). At Ushuaia, measurements were compared to data of the SUV-100 spectroradiometer of the NSF UV monitoring network. This network also included high-latitude sites in Antarctica and the Arctic (Bernhard et al., 2010). Measurements at Ushuaia were discontinued in December 2008. The irradiance scale was traceable to NIST (Bernhard et al., 2003). SUV-100
- 20 measurements at Ushuaia were discontinued in December 2008.

In this study, results of the comparisons done in 2004–2012 were analysed for clear sky conditions (Tables 1 and 2). Results prior to 2004 were analysed and discussed in Lakkala et al. (2005). All comparisons showed that the difference between the traveling reference and the spectroradiometers was within 5% until February 2008. In 2009, the comparison against the Finnish Brewer spectroradiometer in Jokioinen showed a difference of 20% in the erythemally weighted dose rates. In 2010, the NILU-UV reference radiometer was sent to Norway to be recalibrated. The next comparisons were in 2011 and 2012, when the differences were 8% and 26%, correspondingly, which suspect an increased rate of degradation of the instrument. After 2011, the channels of the traveling reference had drifted so severely that even recalibration would not have helped, and measurements with that instrument were not possible after 2012.

**Table 1.** The ratios of erythemally-weighted UV dose rates between the reference NILU-UV and the Brewer spectroradiometers #037 and #107 during 2004–2012.

Date	SZA	Brewer#107/NILU-UV	Brewer#037/NILU-UV
16.7.04	41	1.04	
18.8.05	48	1.01	
10.6.07	44		1.03
16.6.07	37	1.02	
31.8.08	54	1.04	
6.8.09	45	0.80	
17.8.11	48	1.08	
19.7.12	48	1.26	

**Table 2.** The ratios of erythemally-weighted UV dose rates between the reference NILU-UV and the SUV-spectroradiometer during 2004–2006. For 2007–2008 the comparisons were done using the NILU-UV of Ushuaia.

Date	SZA	SUV/NILU-UV
16.1.04	52	0.95
10.12.04	40	0.96
1.3.05	48	1.03
28.11.05	38	0.96
3.3.06	48	0.97
15.11.06	40	0.99
12.2.07	44	1.00
26.10.07	44	1.02
2.2.08	41	1.00

#### 5 2.3.2 Solar comparisons to transfer the irradiance scale

10

The traveling reference typically visited Marambio twice and Ushuaia three times during spring–autumn. During these site visits, solar comparisons were performed to transfer the irradiance scale. Each channel of the site radiometer was scaled to the corresponding channel of the traveling reference. Thus, the calibration of the traveling reference could be used. The method is described in details in Lakkala et al. (2005). In the method, the calibration of the site instruments relevant for solar data was entirely based on the comparisons with the reference instrument<del>and not the lamp scans described earlier</del>.

The sealing coefficient differences between the channels of the traveling reference and the site instrument before the calibration were used as scaling coefficients and their time series is shown in Fig. 3. The differences between the channels 1–3 of the traveling reference and the site NILU-UV were less or around 50% until 2010 for both instruments. The channel 2-two of the instrument of Marambio had however smaller differences than the corresponding channel of the instrument of Ushuaia. The channel four, which degraded the fastest, had more than 80% of difference already in 2008.

- Until 2008For the years 2000–2008 and 2010, all measurements at Ushuaia and Marambio were scaled according to the method by (Lakkala et al., 2005) and the erythemally weighted UV dose rates were calculated. The highest step between calibrations occured in Marambio in January 2000 and it was 11% for erythemally weighted UV dose rates. A typical step between calibrations was from 1 to 5% during summer months. Using the a step of 5% as uncertainty related to the drift
- of the calibration of the site instrument between calibrations, the drift of the traveling reference between calibrations (3% for 2004–2006), the uncertainty of the calibration of the traveling reference (Lakkala et al., 2005), the uncertainties related to the transfer of the irradiance scale (Lakkala et al., 2005) and the uncertainty of the method to calculated the dose rates (Dahlback, 1996), the combined uncertainty was calculated to be 9.5%. The expanded uncertainty was 19% using a coverage factor of 2 (Table 3). In 2009, following the results of solar comparison between the FMI Brewer spectroradiometer and the
- 25 travelling reference, which showed a difference of 20%, all dose rates referenced to the travelling reference were scaled by 20% to account for the drift of this system. In 2010, the reference radiometer was recalibrated by the manufacturer, and the measurements at Ushuaia and Marambio were adjusted to match. Thus the uncertainty of the measurements of this radiometer. The reference radiometer was again compared with the FMI Brewer spectroradiometer in 2011 and UV dose rates measured by the two systems agreed to within 8% the year 2009 was higher than for the other years. The number of days with good measurement accurate measurements that were used in this study, are shown in Table 4 for each year.

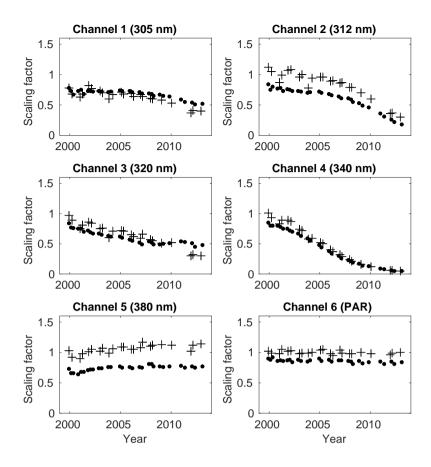


Figure 3. Scaling factors of the UV channels of the Ushuaia (.) and Marambio (+) NILU-UV instruments.

Table 3. Uncertainty budget of the Erythemally weighted UV irradiances during 2000–2008.

Source of uncertainty	Relative uncertainty/ %
Uncertainty of the calibration of the travelling reference provided by the manufacturer	<u>5 (Lakkala et al., 2005)</u>
Drift of the calibration of the reference instrument between calibration	3
Uncertainty of the transfer of the scale of irradiance	2.4 (Lakkala et al., 2005)
Drift of the calibration of the site-instrument between calibrations	5
Uncertainty of calculating the erythemal irradiance	<u>5 (Dahlback, 1996)</u>
Combined uncertainty	<u>9.5</u>
Expanded uncertainty	<u>19</u>

Year	Ushuaia	Marambio
2000	326	294
2001	327	327
2002	311	306
2003	289	317
2004	355	345
2005	338	338
2006	326	325
2007	330	346
2008	342	343
2009	339	335
2010	340	340

Table 4. Number of good measurement days during 2000-20102000-2010.

#### 5 2.4 Potential vorticity analysis

Both Ushuaia and Marambio can be located inside, at the edge or outside the polar vortex. The impact of the location with respect to the polar vortex was studied by calculating the potential vorticity (PV) at Ushuaia and Marambio for measurement days. The modified PV, in which the PV was scaled to 475 K (Lait, 1994) using the ERA interim data, was used. If the PV was higher than the value 36 (in South hemisphere lower than -36), the station could be classified to be inside the vortex.

#### 10 3 Results and Discussion

After the quality assurance procedures mentioned in the earlier sections, the time series of erythemally weighted UV irradiances for Ushuaia and Marambio were calculated for the time period 2000-2010. Maximum daily UV indices and daily doses calculated from the results are presented in Figs. 4 and 5 and discussed in the next sections. The data is archived and can be downloaded at the FMI Arctic Space Centre open access database http://litdb.fmi.fi.

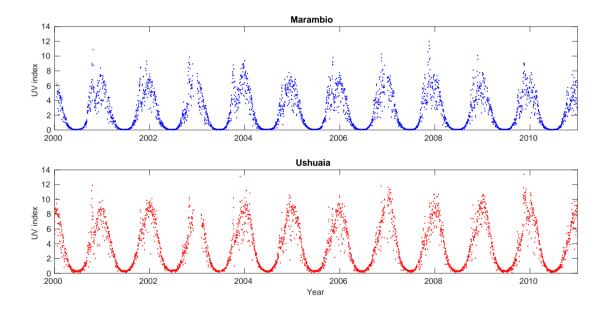


Figure 4. Maximum Daily maximum UV index measured at Marambio and Ushuaia 2000–2010.

#### 3.1 UV index time series

At both stations, the highest UV indices were measured between the end of October and the end of November. As this period is before the summer solstice, factors other than the SZA must be responsible for the observed time-shift of the maximum UV index: The well-known Antarctic ozone depletion occurs during October-November (WMO, 2003) and the snow that still

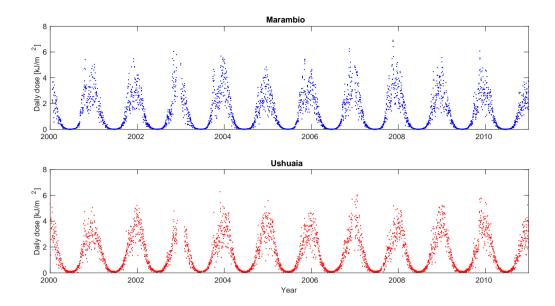


Figure 5. Erythemally weighted daily doses measured at Marambio and Ushuaia 2000–2010.

5 might be on the ground during these months. Both factors increase UV irradiances near the ground. Furthermore, clear skies or broken cloudiness conditions occurred during days when the highest UV values were measured.

At Marambio, extreme (i.e., UV indices of 11 or more, WMO (1997)) daily maximum UV indices were measured on 5 days, and the peak value of 12 was measured on 18 Nov 2007. The daily variation of the UV index on that record day is shown in Fig. 6. The total ozone amount on that day was 213 DU. The daily maximum UV was very high (i.e., 8-10) on 85 days and high (i.e, 6-7) on 524 days. The daily elimatological means values of the daily climatology calculated from the daily maximum UV indices (2000–2010) were the highest, UV index of 6, in late November, December and January (Fig. 7). Fig. 7 shows that the daily elimatological means values of the daily climatology were not symmetrically distributed around mid summer, but increased in October compared to days with the same noon-time SZA in autumn. Peak UV indices occurred between late October and late November, depending on how long the polar vortex stayed above the station. That was also the period

15 during which the maximum UV indices deviated the most from the daily means climatology (Fig. 8). Also snow typically enhanced UV radiation in spring (Meinander et al., 2014). The RAF was calculated to be 1.16, when excluding the winter months May-July from the analysis. When largest ozone decreases occured, the RAF could exceed 1.25 (Fig. 9) suggesting an additional increase due to snow on the ground. Measured mid summer clear sky daily maximum UV indices were between 7 and 8. In June, no UV radiation was measured at the station. For comparison: At the same latitude in Nothern hemisphere, in Finland, the highest measured UV indices were between 5 and 6. The main reasons for higher summer UV indices in Marambio were lower total ozone amounts and shorter smaller Earth-Sun distance.

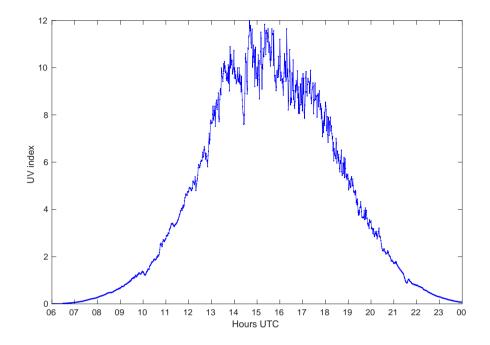
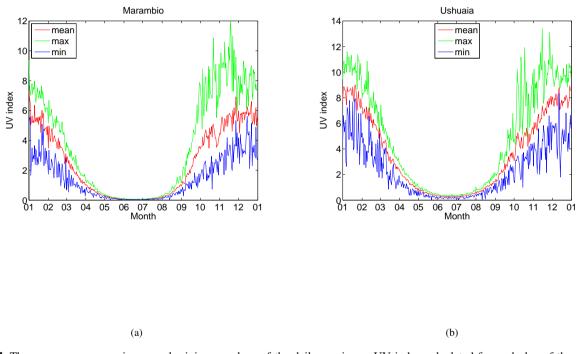
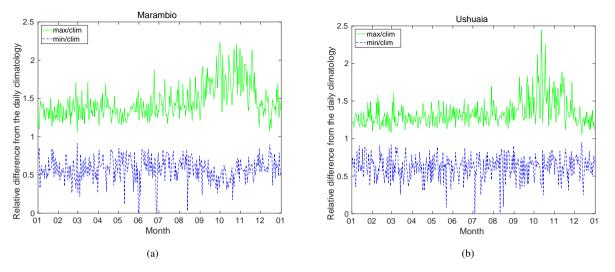


Figure 6. UV indices measured by the NILU-UV radiometer at Marambio on the day of record high UV index on 18 November 2007.

- At Ushuaia, extreme daily maximum UV index was measured on 35 days. The highest value, UV index 13, was measured twice: on 27 Nov 2003 and on 15 Nov 2009. The corresponding total ozone amount of 27 Nov 2003 was 268 DU, but no ozone data was recorded for 15 Nov 2009. However, low ozone was evident as the total ozone amounts of the previous and next days were 176 and 215 DU, respectively. Very high daily maximum UV index was measured on 515 days and high on 1045 days. Daily <u>elimatological means climatology</u> calculated from the daily maximum UV indices 2000–2010 were highest, UV index
- 10 8–9, in January (Fig. 7). The graph of the daily <u>elimatological mean-climatology</u> showed the same, but not as pronounced, features than at Marambio: Enhanced daily maximum UV indices in October (Fig. 8). When the station was inside the polar vortex in October, the maximum UV indices could be as high as those measured in the summer, when they were between 10 and 11.
- The deviations from the daily elimatological mean-climatology were calculated for each day (Fig. 10) and analyzed with respect to the location of the polar vortex. All deviations larger than 80% from the mean occurred when the stations were inside the polar vortex. At Ushuaia this was the case especially in 2000–2004 and in 2009. At Marambio, deviations of over 100% were measured in 2000, 2002, 2003 and 2007. Highest deviations were 122% on 1 Oct 2003 and 145% on 12 Oct 2000 at Marambio and Ushuaia, respectively. During cloudy days, the daily maximum UV index did not reached reach high values, even if the station was inside the polar vortex. Lowest values occured during days with heavy cloudiness or rainy days, with deviations of 100% from the daily meanclimatology.



**Figure 7.** The averagemean, maximum and minimum values of the daily maximum UV index calculated for each day of the year in the period 2000–2010 in a) Marambio b) Ushuaia.



**Figure 8.** The maximum and minimum values of the daily maximum UV index divided by the <u>mean daily climatology</u> of maximum UV index on that day in a) Marambio b) Ushuaia 2000-2010.

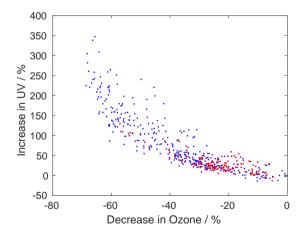
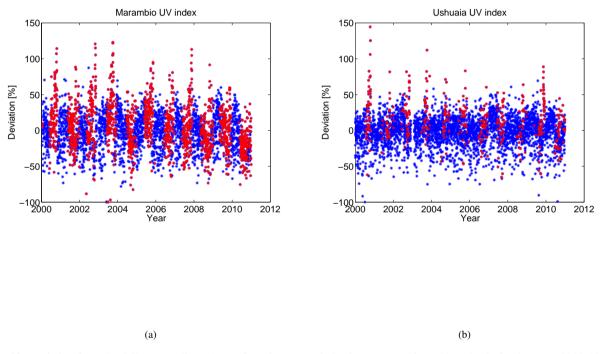


Figure 9. Total ozone decrease from the highest total ozone value measured in Marambio during the measurement period 2000–2010 and the corresponding increase in erythemally weighted irradiances. Blue dots corresponds to months when there is probably more snow on the ground (August, September, October, March and April) and the red dots correspond to summer months (November, December, January and February).



**Figure 10.** Deviation from the daily <u>mean-climatology of maximum UV</u> index in a) Marambio and b)Ushuaia for the years 2000-2010. Days for which the station was inside the polar vortex are indicated with a red symbols, the others in blue.

#### 3.2 UV daily dose time series 5

Daily doses were strongly dependent on the total ozone amount and cloudiness conditions. For some days with low total ozone and clear sky, the daily dose time series graph showed that peak daily doses were more pronounced than peak UV indices (Fig. 4). The highest daily doses were around  $6 kJ/m^2$ , at both Marambio and Ushuaia , even though the difference in latitude between these sites is 10°. At Marambio, a daily dose of 7  $kJ/m^2$  was measured in 2007, when the station was inside the polar

vortex and episodes with low total ozone occurred. The highest daily doses were measured at Ushuaia and Marambio on 27 10 Nov 2003 and 19 Nov 2007, respectively.

The graph of the daily elimatological mean climatology of the daily doses (FigsFig. 11and 12) showed features similar to those of the daily elimatological mean climatology of maximum UV indices (Fig. 7). The impact of low ozone episodes in October is seen, but not as clearly as for daily maximum UV indices. At Marambio, the highest mean daily doses occurred

- 15 between late November and the end of December. When the station was inside the polar vortex and clear sky conditions occurred, the typical daily dose was around 5  $kJ/m^2$  in November. The typical daily dose was between 2.5 and 3.5  $kJ/m^2$  in summer, but around 4.5  $kJ/m^2$  on sunny days. The high fluctuation in maximum values (Fig. 12) is due to the impact of year to year changes in cloudiness and total ozone.
- At Ushuaia, typically daily doses were between 3 and 4  $kJ/m^2$  in summer. The graph of the elimatological daily mean daily climatology of daily doses showed maxima between late December and the end of January. Highest daily doses, between 4.5 20 and 5  $kJ/m^2$ , occurred typically in late January. However, daily doses of the same amount were measured on days with clear sky and low total ozone in late November. Compared to Marambio, the period of highest daily doses was shorter. At Marambio the period of highest daily doses, over 4  $kJ/m^2$ , started already in the beginning of October, while it started at Ushuaia in November. At Marambio, in addition to more frequent low total ozone conditions, also the snow on the ground increased the UV irradiances due to high surface reflection of the incoming radiation. 25

Both stations were inside the polar vortex on most of the days when the daily dose deviated by more than 75% from the daily elimatological mean climatology (Fig. 13). At Ushuaia this was the case in 2000 – 2003. The highest deviation of 141% from the mean was measured on 6 Oct 2003. At Marambio, in all cases except one, the station was inside the polar vortex when the daily doses deviated by more than 100% from the elimatological mean daily climatology of the day. The highest deviation of 150% from the daily elimatological mean-climatology was measured on 3 Nov 2002. Deviations exceeded 100% from the daily mean-climatology also in 2000, 2003, 2005, 2007 and 2008.

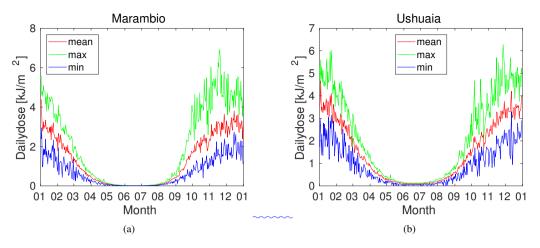
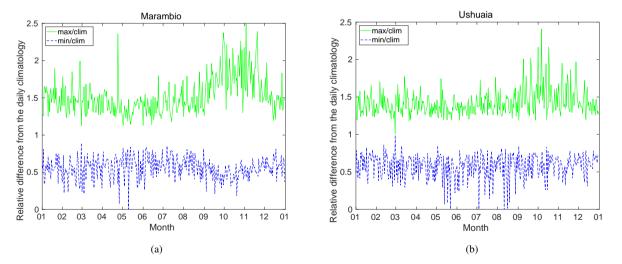
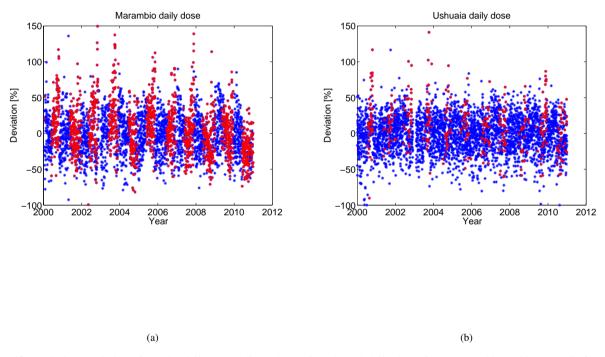


Figure 11. The mean, maximum and minimum ealeulated for each day of the year from the erythemally weighted dailydoses daily dose  $[J/m^2kJ/m^2]$  measured calculated from the respective days of each year of observations in a) Marambio b) Ushuaia 2000-2010.



**Figure 12.** The maximum and minimum erythemally weighted daily doses calculated for each day from respective days of the each year from the erythemally weighted dailydoses of observations divided by the mean of the same day daily climatology in a) Marambio b) Ushuaia 2000-2010.



**Figure 13.** Deviation Deviations from the daily mean climatology of erythemal daily doses in a) Marambio and b) Ushuaia for the years 2000-2010. Days for which the station was inside the polar vortex are indicated with a red symbol, the others in blue.

#### 5 4 Future

After the operation of the Antarctic NILU-UV network until 2014, co-operative Argentinian-Finnish UV measurements were continued using erythemally weighted UV sensors. In cooperation between FMI and Servicio Meteorológico Nacional (SMN), Argentina, UV broadband radiometers, type SL501A manufactured by Solar Light Co., were installed in <u>Marambio in 4</u> February 2013 to measure the local bihemispherical UV albedo (Meinander et al., 2014). One instrument was installed to

- 10 measure UV radiation reflected from the ground, and a second one to measure global UV irradiance. Using these measurements, the UV index time series could be continued, as those measurement are also traceable, via the primary calibration lamp, to the National Standard Laboratory MIKES, Aalto University, Finland. The official SL501A trace is to the primary standard of the National Institute of Standards and Technology (NIST). The difference in calibration coefficients using NIST and MIKES has been found to be less than 2%, and in comparison of spectral UV irradiance scales maintained by NIST, PTB and Aalto
- 15 University (HUT), no major differences have been found (Jokela et al., 2000). Hence, the irradiance scale of these new SL501A radiometers is comparable to that of the NILU-UV measurements within 2%.

Likewise in the scientific cooperation between SMN and FMI, a new GUV-2511 multi-filter radiometer, manufactured by Biospherical Instruments Inc., was installed at Marambio in 2017. It features five channels at UV wavelengths and two channels at visible radiation wavelengths (Bernhard et al., 2005). Two GUV radiometers were bought by FMI to ensure continuous

- 20 measurements at Marambio: while one of them is installed at the site, the other one is calibrated and also available for comparison campaigns. These measurements continue the NILU-UV time series, as the central wavelengths of the channels are nearly identical to those of the NILU-UV channels and the bandwidths are similar. The first calibration of the GUVs were performed by Biospherical Instruments, Inc. Comparisons with FMI's spectroradiometers showed that the measurements were within 5% from each other. These results are promising for the future of the UV time series at Marambio, and for monitoring the impact
- 25 of the predicted ozone layer recovery on surface UV radiation in the Antarctic. The measurements will also serve as important validation data for satellite measurements, similar to the the NILU-UV measurements previously.

#### 5 Conclusions

30

In this paper, UV measurements from the Antarctic NILU-UV network at Marambio (64°S) and Ushuaia (54°S) were quality controlled and analyzed. The studied time series included measurements from 2000 to 2013. The quality assurance procedures included in situ lamp measurements and solar comparisons with the traveling reference of the network.

The results showed that after the year 2010 the filters of the radiometers degraded and their sensitivity to UV radiation was reduced by more than 80% in some channels. Due to this problem, accurate data were available only until 2010. The irradiances measured by the traveling reference were compared to those of well maintained spectroradiometers in Finland and Argentina. From these results, also The results showed that the traveling reference data were found to be accurate only until 2010. The data of the stations of Ushuaia and Marambio were scaled to that of the traveling reference and homogenized can be used to transfer the calibration to the stations and that the uncertainty of the measurements increased after 2008. Homogenized UV irradiance time series were calculated.

- In both Ushuaia's and Marambio's time series, the effect of low total ozone amount during days that the station was inside the polar vortex was clearly seen. At Marambio, daily doses for late October and November could even exceed those measured in the summer. The maximum daily dose was as large as  $7 kJ/m^2$  in November 2007. Highest daily doses were typically around  $6 kJ/m^2$  at both stations, even if the difference in latitude is 10°. In the whole measurement period (2000–2010) there were 5 days when extreme UV indices and 85 days when very high UV indices were measured at Marambio. The highest daily
- 10 maximum UV index of 12 was measured in November 2007. At Ushuaia, extreme UV indices were measured on 35 days. The highest UV index of 13 was measured twice: in November 2003 and 2009. Largest deviations from the daily means were over 100% and occured in October and November when the stations were inside the polar vortex. At Ushuaia, the largest deviations from the mean occured during the first years of the time series, i.e, 2000–2003.

Code availability. TEXT

15 Data availability. The data can be downloaded at the FMI Arctic Space Centre open access database http://litdb.fmi.fi.

Author contributions. K. Lakkala: Primarily responsible for the QA of the UV data, analysed the data and led the manuscript preparation.

- A. Redondas: Principal Investigator on the Antarctic NILU-UV project and contributed to the manuscript.
- O. Meinander: QC of the traveling reference. One of the main authors of the FARPOCC and SAARA research plans of PI prof. Esko Kyrö.
- L. Thölix: Analyzed the polar vortex and contributed to the manuscript.

- F. Almansa: QA of NILU-UV Antarctic network.
- V. Carreno: QA of NILU-UV Antarctic network.
- R. D. Garcia: QA of NILU-UV Antarctic network.
- C. Torres: QA of NILU-UV Antarctic network.
- 25 G. Deferrari: Daily housekeeping and QC of NILU-UV data at Ushuaia.
  - H. Ochoa: Daily housekeeping and QC of NILU-UV data at Marambio.

G. Bernhard: QA of the SUV spectroradiometer data at Ushuaia. Calibration of FMI's GUV radiometers. Contributed to the writing of the manuscript

R. Sanchez: QC of the GUV measurements at Marambio.

G. de Leeuw: Overseeing the work in the group and contributed to the writing of the manuscript.

Competing interests. No competing interests are present.

<sup>20</sup> B. Hamari: QA of the UV data 2007-2013.

5 Acknowledgements. The Academy of Finland has given financial support for this work through projects FARPOCC and SAARA. The MAR Project was financed by the National R+D Plan of the Spanish Ministry of Science and Technology (National Research Program in the Antarctic) under contract REN2000-0245-C02-01. The SUV-100 spectroradiometer UV data from Ushuaia were provided by the NSF UV Monitoring Network, operated by Biospherical Instruments Inc. and funded by the U.S. National Science Foundation's Office of Polar Programs. We thank the operators of the MAR project at Marambio and Ushuaia. Arne Dahlback and Björn Johnsen are acknowledged for the calibration of the NILU-UVs. We thank the operators of the FMI Brewers at Sodankylä and Jokioinen. Tapani Koskela is acknowledged for processing the Brewer data of Jokioinen. The OMI/AURA and TOMS teams are thanked for the satellite ozone data. We thank Eija Asmi and Edith Rodriguez from FMI for logistics, and Lasse Ylianttila from the Finnish Radiation and Nucleation Safety Authority (STUK) for calibration of SL501A radiometers.

#### 5 References

- Andrady, A., Aucamp, P. J., Bais, A. F., Ballaré, C. L., Björn, L. O., Borman, J. F., Calwell, M., Cullen, A. P., Erickson, D. J., De Gruil, F. R., Häder, D.-F., Ilyas, M., Kulandaivelu, G., Kumar, H. D., Longstreth, J., McKenzie, R. L., Norval, M., Paul, N., Redhwi, H. H., Smith, R. C., Salomo, K. R., Sulzberger, B., Takizawa, Y., Tang, X., Teramura, A. H., Torikai, A., Van Der Leun, J. C., Wilson, S. R., Worrest, R. C., and Zepp, R.: United Nations Environment Programme, Environmental Effects Assessment Panel, 2009, Envi-
- 10 ronmental Effects of ozone depletion and its interactions with climate change: progress report, 2008., Photoch. Photobio. Sci., 8, 13–22, https://doi.org/10.1039/b820432m, 2009.

Bernhard, G., Booth, C. R., and McPeters, R. D.: Calculation of total column ozone from global UV spectra at high latitudes, J. Geophys. Res., 108, https://doi.org/10.1029/2003JD003450, 2003.

Bernhard, G., Booths, C. R., and Ehramjian, J. C.: Real-time ultraviolet and column ozone from multichannel ultraviolet radiometers de-

- 15 ployed in the National Science Foundation's ultraviolet monitoring network, Opt. Eng., 44, 041 011, https://doi.org/10.1117/1.1887195, 2005.
  - Bernhard, G., Booth, C. R., and Ehramjian, J. C.: Comparison of UV irradiance measurements at Summit, Greenland; Barrow, Alaska; and South Pole, Antarctica, Atmospheric Chemistry and Physics, 8, 4799–4810, https://doi.org/10.5194/acp-8-4799-2008, https://www. atmos-chem-phys.net/8/4799/2008/, 2008.
- 20 Bernhard, G., Booth, C. R., and Ehramjian, J. C.: Climatology of Ultraviolet Radiation at High Latitudes Derived from Measurements of the National Science Foundation's Ultraviolet Spectral Irradiance Monitoring Network, In UV Radiation in Global Climate Change: Measurements, Modeling and Effects on Ecosystems, edited by W. Gao, D. L. Schmoldt, and J. R. Slusser, Tsinghua University Press, Beijing and Springer, New York, 2010.

Booth, C., Lucas, T., Morrow, J., Weiler, C., and Penhale, P.: The United States National Science Foundation's Polar Network for Monitoring

- 25 Ultraviolet Radiation, in: Antarctic Research Series, edited by Weiler, C. and Penhale, P., vol. 62, pp. 17–37, American Geophysical Union, Washington, D.C., 1994.
  - Booth, C. R. and Madronich, S.: Radiation amplification factors: improved formulation accounts for large increases in ultraviolet radiation associated with Antarctic ozone depletion, Ultraviolet radiation in Antractica: Measurements and biological effects, Antarctic Research Series, 1994.
- 30 Dahlback, A.: Measurements of biologically effective UV doses, total ozone abundances, and cloud effects with multichannel, moderate bandwidth filter instruments, Appl. Opt., 35, 6514–6521, 1996.
  - Eleftheratos, K., Kazadzis, S., Zerefos, C. S., Tourpali, K., Meleti, C., Balis, D., Zyrichidou, I., Lakkala, K., Feister, U., Koskela, T., Heikkilä, A., and Karhu, J. M.: Ozone and Spectroradiometric UV Changes in the Past 20 Years over High Latitudes, Atmosphere-Ocean, 53, 117– 125, https://doi.org/10.1080/07055900.2014.919897, https://doi.org/10.1080/07055900.2014.919897, 2015.
- 35 Farman, J., Gardiner, B., and Shanklin, J.: Large losses of total ozone in Antarctica reveal seasonal CLO<sub>x</sub>/NO<sub>x</sub> interaction, Nature, 315, 207–210, 1985.
  - Heikkilä, A., Mäkelä, J., Lakkala, K., Meinander, O., Kaurola, J., Koskela, T., Karhu, J. M., Karppinen, T., Kyrö, E., and de Leeuw, G.: In search of traceability: Two decades of calibrated Brewer UV measurements in Sodankylä and Jokioinen, Geosci. Instrum. Method. Data Syst., 5, 531, 2016.

- 5 Høiskar, B., Haugen, R., Danielsen, T., Kylling, A., Edvardsen, K., Dahlback, A., Johnsen, B., Blumthaler, M., and Schreder, J.: Multichannel moderate-bandwidth filter instrument for measurement of the ozone-column anount, cloud transmittance, and ultraviolet dose rates, Appl. Optics, 42, 3472–3479, 2003.
  - Johnsen, B., Mikkelborg, O., Hannevik, M., Nilsen, L., Saxebøl, G., and Blaasaas, K.: The Norwegian UV-monitoring program Period 1995/96 to 2001, Strålevern Rapport 2002:4, Norwegian Radiation Protection Authority, Østerås, 2002.
- 10 Jokela, K., Ylianttila, L., Visuri, R., Leszczynski, K., Kärhä, P., and Ikonen, E.: Intercomparison of lamp and detector-based UV-irradiance scales for solar UV radiometry, Journal of Geophysical Research: Atmospheres, 105, 4821–4827, https://doi.org/10.1029/1999JD900398, http://dx.doi.org/10.1029/1999JD900398, 2000.
  - Karhu, J. A., Taalas, P., Damski, J., Kaurola, J., Ginzburg, M., Villanueva, C. A., Piacentini, E., and Garcia, M.: Vertical distribution of ozone at Marambio, Antarctic Peninsula, during 1987–1999, Journal of Geophysical Research: Atmospheres, 108, n/a–n/a, https://doi.org/10.1029/2003JD001435, http://dx.doi.org/10.1029/2003JD001435, 4545, 2003.
  - Lait, L. R.: An Alternative Form for Potential Vorticity, Journal of the Atmospheric Sciences, 51, 1754–1759, https://doi.org/10.1175/1520-0469(1994)051<1754:AAFFPV>2.0.CO;2, https://doi.org/10.1175/1520-0469(1994)051<1754:AAFFPV>2.0.CO;2, 1994.

15

20

35

- Lakkala, K., Redondas, A., Meinander, O., Torres, C., Koskela, T., Cuevas, E., Taalas, P., Dahlback, A., Deferrari, G., Edvardsen, K., and Ochoa, H.: Quality assurance of the solar UV network in the Antarctic, J. Geophys. Res., 110, https://doi.org/10.1029/2004JD005584, 2005.
- Lakkala, K., Arola, A., Heikkilä, A., Kaurola, J., Koskela, T., Kyrö, E., Lindfors, A., Meinander, O., Tanskanen, A., Gröbner, J., and Hülsen, G.: Quality assurance of the Brewer spectral UV measurements in Finland, Atmos. Chem. Phys., 8, 3369–3383, 2008.

McKinlay, A. F. and Diffey, B. L.: A reference action spectrum for ultraviolet induced erythema in human skin, CIE J., 6, 17–22, 1987.

- Meinander, O., Asmi, E., Neitola, K., Laurila, T., Lihavainen, H., Aaltonen, V., Poikonen, A., Aarva, A., Rodriguez, E., Kyrö, E., de Leeuw,
- 25 G., Sanchez, R., and Mei, M.: Finnish-Argentinian Measurements on Snow UV Albedo in Marambio, In: FINNARP. Science and Support in Antarctica. Finnish Antarctic Research Program (FINNARP), ISBN 978-951-697-841-6, 2014., 2014.
  - Pazmiño, A., Godin-Beekmann, S., Ginzburg, M., Bekki, S., Hauchecorne, A., Piacentini, R., and Quel, E.: Impact of Antarctic polar vortex occurences on total ozone and UVB radiation at southern Argentinean and Antarctic stations during 1997–2003 period, J. Geophys. Res., 110, https://doi.org/10.1029/2004JD005304, 2005.
- 30 Solomon, S., Ivy, D. J., Kinnison, D., Mills, M. J., Neely, R. R., and Schmidt, A.: Emergence of healing in the Antarctic ozone layer, Science, https://doi.org/10.1126/science.aae0061, http://science.sciencemag.org/content/early/2016/06/30/science.aae0061, 2016.
  - WMO: Scientific Assessment of Stratospheric ozone: 1989, World Meteorological Organization (WMO), Global Ozone Research and Monitoring Project–Report No. 20, 1990.
  - WMO: Report of the WMO-WHO Meeting of Experts on Standardization of UV Indices and their Dissemination to the Public, World Meteorological Organization (WMO), Global Atmosphere Watch Report No. 127, 1997.
  - WMO: Scientific Assessment of Ozone Depletion:1998, World Meteorological Organization (WMO), Global Ozone Research and Monitoring Project–Report No. 44, 1999.
- 400 WMO: Scientific Assessment of Ozone Depletion:2002, World Meteorological Organization (WMO), Global Ozone Research and Monitoring Project–Report No. 47, 2003.
  - WMO: Scientific Assessment of Ozone Depletion: 2006, World Meteorological Organization (WMO), Global Ozone Research and Monitoring Project–Report No. 50, 2007.