



- Analysis of European ozone trends in the period 1995–2014
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8 Abstract

9 Surface-based measurements from the EMEP network are used to estimate the changes in surface ozone levels during the 1995-2014 period over Europe. It is shown that a significantly decreasing 10 trend in the 95th percentile ozone concentrations has occurred, especially during noontime (0.9 11 $\mu g/m^3/y$), while the 5th percentile ozone concentrations continued to increase with a trend of 0.3 12 $\mu g/m^3/y$ during the study period. With the help of numerical simulations performed with the 13 global chemistry-climate model EMAC, the importance of anthropogenic emissions changes in 14 15 determining these changes are investigated. The EMAC model is found to successfully capture the observed temporal variability in mean ozone concentrations, as well as the contrast in the 16 trends of 95th and 5th percentile ozone over Europe. Sensitivity simulations and statistical analysis 17 show that a decrease in European anthropogenic emissions had contrasting effects on surface 18 ozone trends between the 95th and 5th percentile levels, and that background ozone levels have 19 been influenced by hemispheric transport, while climate variability generally regulated the inter-20 annual variations of surface ozone in Europe. 21

22 1. Introduction

Tropospheric ozone has detrimental effects on human health, and elevated concentrations at the 23 24 surface are of concern over most of the European region (Hjellbrekke and Solberg, 2002; WHO, 2013; EEA, 2013; Lelieveld et al., 2015). The European Union (EU) Air Quality Directive sets 25 four standards for surface ozone to reduce its impacts on human health and crop yields 26 (http://eur-lex.europa.eu/legal-content/EN/TXT/HTML/?uri=CELEX:32008L0050&from=EN). 27 The standards are: information threshold, (1-hour average: 180 µg/m³), alert threshold (1-hour 28 average: 240 μ g/m³), long-term objective (maximum diurnal 8-hour mean: 120 μ g/m³), and the 29 target value (long-term objective should not be exceeded more than 25 days per year, averaged 30 over 3 years). Exceedances are particularly frequent in regions close to high ozone precursor 31 emissions during summer with stagnant meteorological conditions, associated with persistent 32 high temperatures. Since a substantial decrease in precursor concentrations has been achieved in 33

34 Europe in recent decades, the number of exceedances has declined (Guerreiro et al., 2014), in line





35 with a long-term downward trend of pollution emissions (Colette et al., 2011; Wilson et al., 2012).

Further, a number of studies have shown that European ozone levels are on average decreasing in the last 20 years (as example, Jonson et al, 2010). Nevertheless, background ozone changes over

Europe are not so clear (Wilson et al., 2012), being sensitive to climate conditions and

- intercontinental transport of O_3 and its precursors, and are significant in view of tropospheric chemistry (Lelieveld and Dentener, 2000).
- The response of surface ozone to a changing climate, with potentially more frequent heat 41 42 extremes (Bloomer et al., 2009; Jacob and Winner, 2009; Cooper et al., 2012; Fu et al., 2015; Lin et al., 2015; Simon et al., 2015), and concurrent changes in anthropogenic emissions of precursor 43 gases (Bloomer et al., 2009; Fu et al., 2015; Strode et al., 2015; Yan et al., 2017) may pose a 44 45 challenge for air quality management. Observation and model-based analyses of ozone trends in 46 responses to climate change (Bloomer et al., 2009), precursor emissions (Bloomer et al., 2009; Lefohn et al., 2010), and long-range transport (Lin et al., 2015) have been conducted for North 47 America (Strode et al., 2015; Lin et al., 2017; Yan et al., 2017), several Asian regions (Brown-48 Steiner et al., 2015; Lin et al., 2017) and also for Europe (Meleux et al., 2007, Wilson et al., 2012, 49 Jonson et al., 2006). For Europe, the connection between climate and ozone levels has been 50 subject of large number of studies, notably to investigate the effects of climate change on surface 51 ozone levels (Langner et al., 2005; Meleux et al., 2007; Colette et al., 2011; Langner et al., 2012.) 52

53 Tropospheric ozone is produced photochemically during daytime, mainly from the photolysis of nitrogen dioxide (NO₂), while NO₂ levels are strongly influenced by radicals and their precursors, 54 including organic compounds. Due to the complex photo-chemistry involved, the amount of 55 ozone formed responds nonlinearly to changes in precursor emissions and is sensitive to 56 variations in air temperature, radiation and other climatic factors (Fu et al., 2015; Monks et al., 57 2015; Coates et al., 2016). Ozone can be destroyed via reactions with NO_x (i.e., ozone titration) 58 especially during nighttime, and thus a reduction in NO_x emissions could result in more ozone 59 (Jhun et al., 2014; Yan et al., 2017). Previous studies of European ozone have focused on daytime 60 or diurnal mean ozone with little attention paid to the daytime-nighttime contrast in ozone 61 changes (Colette et al., 2011; Wilson et al., 2012; Guerreiro et al., 2014). 62

Our work contrasts the trends of the monthly 5th and 95th percentile European ozone levels at hourly levels over the period 1995–2014, based on the hourly ozone measurements from the EMEP network. Additionally, numerical simulations from the global chemistry-climate model ECHAM5/MESSy (EMAC) are conducted to evaluate the model's ability in capturing ozone trends over Europe and to investigate the underlying importance of the meteorology and emission changes for the observed ozone trends.

The manuscript is organized as follows: the observational dataset, model simulations and analysis methods are described in Section 2. In Section 3, the average linear trends for the European domain are estimated and analyzed separately for the monthly, seasonal and annual surface 5th, 50th, and 95th percentiles of the observed ozone concentrations. We then compare the observed





73 ozone trends and variability to results of the atmospheric chemistry – general circulation model

EMAC. To investigate the effects of anthropogenic emissions and climate variability on observed

European ozone changes, we conduct a sensitivity simulation with constant emissions and statistical analysis with the ERA-Interim 2-meter temperature data in Section 4. Followed by the

statistical analysis with the ERA-Interim 2conclusions in Section 5.

78 2. Methods and Data

79 **2.1 Ozone measurements**

The hourly ground-level ozone measurements over 1995–2014 have been obtained from the 80 81 Chemical Coordination Centre of European Monitoring and Evaluation Programme (EMEP) network (http://www.nilu.no/projects/ccc/emepdata.html). Table 1 shows the number of 82 measurement sites (varies from 113 to 137) and the percentage of missing hourly data in each 83 year. Fig. 1 further shows the site distribution. Since many of the stations are not operating 84 continuously during the study period (Fig. 1), we have included only those sites in the analysis 85 which fulfill the criteria defined by Cooper et al. (2012). Such data selection criteria are further 86 applied for the US ozone trends analysis with the EPA-AQS measurements by Yan et al. (2017). 87 First, we discard the observational days with the valid hourly data less than 66.7% in any daytime 88 or nighttime. Then, we discard the particular season with less than 60 days containing valid data 89 90 in any season. Finally, for any season, we keep the data with valid seasonal mean ozone more than 15 years during 1995–2014; otherwise we discard the data in all years for the particular 91 season. Fig. 1 shows the final selected 93 sites satisfying above criteria for the analysis. 92

We calculated the linear trends for the European surface ozone at individual hours, and mean values for daytime (local time: 07:00–19:00), nighttime (local time: 19:00–07:00) and full days (24 h). For each daytime or nighttime period, the missing data varies between 6.8 and 34.6% (Table 1). The monthly 5th, 50th and 95th percentile ozone concentrations for each period (per hour, daytime, nighttime and diurnal) are derived from the lowest, middle and highest 5th percentile hourly ozone mixing ratios of the corresponding period in each month.

To calculate the ozone trends per hour, during daytime, nighttime and per day, we then use the
following statistical trend model (Weatherhead et al., 1998; Yoon and Pozzer, 2014):

$$101 \qquad Y_t = \mu + S_t + \omega X_t + N_t$$

Where Y_t denotes the monthly time series of ozone, μ is a constant term representing the offset, $X_t = t/12$ (with t as month) the number of years in the timeseries, and ω is the magnitude of the trend per year. S_t is a seasonal component in the trend estimates. N_t is the residual term of the interpolation. As the seasonal component does not have much impact on the statistical properties of the estimates of the other terms in the model, we use the deseasonalized monthly data to perform the trend analysis with a model of the form:





 $Y_t = \mu + \omega X_t + N_t$

Using this formulation the linear trends are also analyzed separately for the observed monthly, seasonal and annual surface ozone concentration. The 5^{th} , 50^{th} and 95^{th} percentile ozone

110 concentrations are derived from the lowest, middle and highest 5th percentile of the full set of

111 measurements (i.e., hourly ozone mixing ratios) of the corresponding period.

112 The standard deviation of ozone trends over the European stations is calculated with:

113
$$\sigma = \sqrt{\frac{1}{N}\sum_{i=1}^{N}(\omega_i - \alpha)^2}$$

where *N* is the total number of sites, ω_i is ozone trend at individual sites and α represents the average ozone trend.

116 **2.2 ERA-Interim 2-meter temperature data**

The 2-meter temperature data is from the reanalysis product ERA-Interim, provided by the 117 European Centre for Medium Range Weather Forecast (ECMWF) Public Datasets web interface 118 (http://apps.ecmwf.int/datasets/), covering the data-rich period from 1979 and continuing in real 119 time (Dee et al., 2011). Compared to the ERA-40, the ERA-Interim has an improved 120 representation of the hydrological cycle, and stratospheric circulation (Dee and Uppala, 2009; 121 122 Dee et al., 2011). The ERA-Interim atmospheric model and reanalysis system uses cycle 31r2 of 123 ECMWF's Integrated Forecast System (IFS), configured for 60 vertical levels up to 0.1 hPa. The horizontal-spatial resolution is either in a full T255 spectral resolution or in the corresponding 124 N128 reduced Gaussian grid (Dee et al., 2011). ERA-Interim assimilates four analyses per day, at 125 00, 06, 12 and 18 UTC. ECMWF public website provides a large variety of data in uniform 126 lat/long grids varying from 0.125° to 3°. Out of those, here, we analyze the monthly mean 2-127 128 meter temperature data which are archived on the 0.75° latitude by 0.75° longitude grid. Additional information (e.g. on current data availability) is available on the ECMWF website at 129 http://www.ecmwf.int/research/era. 130

131 **2.3 Atmospheric chemistry modeling**

The ECHAM5/MESSy Atmospheric Chemistry (EMAC) model has been used to simulate surface ozone for the 1995–2014 periods. The EMAC model applies the second version of the Modular Earth Submodel System (MESSy2) to link multi-institutional computer codes (Jockel et al., 2016). The core atmospheric model is the 5th generation European Centre Hamburg general circulation model (ECHAM5) (Roeckner et al., 2006). EMAC simulated gas-phase tracers as well as aerosols have been extensively evaluated in previous studies (e. g. Pozzer et al., 2007; Pozzer et al., 2012).





139 In this work, we use the archived RC1SD-base-10a simulation results from the EMAC model conducted by the ESCiMo project (Jockel et al., 2016). The model results were simulated with 140 version 5.3.02 for ECHAM5 and version 2.51 for MESSy. The archived data were obtained with 141 a T42L90MA spatial resolution, i.e., with a T42 spherical representation which is corresponding 142 to a quadratic Gaussian grid with approximately 2.8 latitude by 2.8 longitude, and 90 levels in the 143 vertical, with the top level up to 0.01 hPa. To reproduce the observed meteorology, the method of 144 Newtonian relaxation towards ERA-Interim reanalysis data (Dee et al., 2011) is applied to 145 weakly nudge the dynamics of the general circulation model. Differently from the work of J ckel 146 147 et al. (2016), the model was re-run to cover the full period of measurements and also with a 1hourly temporal resolution for ozone, in order to compare model results with hourly 148 observational data. We also conducted a sensitivity simulation in which the anthropogenic 149 150 emissions were kept constant throughout the years to investigate the effects of emissions on ozone trends. 151

Anthropogenic emissions in the model are incorporated as prescribed sources following the Chemistry-Climate Model Initiative (CCMI) recommendations (Eyring et al., 2013), using the MACCity (Monitoring Atmospheric Composition & Climate/City Zero Energy) emission inventory, which includes a seasonal cycle (monthly resolved). Additionally, the emissions are vertically distributed as described by Pozzer et al. (2009). Since the total NMVOCs (non-methane volatile organic compounds) values for anthropogenic sectors are not provided by the MACCity raw dataset, they are recalculated from the corresponding species (Jockel et al., 2016).

Emissions from natural sources have been prescribed as well, either as monthly resolved or annually constant climatology. The spatial and temporal distributions of biogenic NMHCs are based on Global Emissions InitiAtive (GEIA). Further details of the model setup on the emissions, physical and chemical processes as well as the model evaluation with observations can be found in J öckel et al. (2016).

164 **3. Results**

165 **3.1 Ozone trends in EMEP measurements**

Fig. 2 shows the trends in ozone concentrations (monthly mean, 5th, 50th and 95th percentile) over 166 Europe during the 1995-2014 period, for each hour of the day. While the average ozone 167 concentrations (and 50th percentiles) do not show significant trends, the 5th and 95th percentile 168 ozone show significant trends with a clear diel cycle. The 95th percentile ozone shows a 169 decreasing trend over Europe during the 1995-2014 period with the trend being most pronounced 170 $(-0.9 \pm 0.5 \ \mu g/m^3/y)$ during midday (1100-1500 h). 95th percentile ozone concentrations also 171 show a decreasing trend during the night, however the trends are observed to be smaller (-0.5 \pm 172 0.35 $\mu g/m^3/y$). Here the standard deviation depicts the variability of the trends among the stations, 173 and therefore reflects the almost homogeneous decrease over entire Europe. Interestingly, in 174 contrast with the 95th percentile, the 5th percentile ozone over Europe shows an increasing trend 175





especially during midday (0.3 \pm 0.16 µg/m³/y). Further, the temporal evolutions of ozone 176 anomalies during the 1995–2014 period are shown for 5th and 95th percentile in Fig. S1. The 95th 177 percentile ozone trend indicates a general decline in the photochemical buildup of ozone during 178 noon hours, with the exception of strongly enhanced ozone during 2003. The inter-annual 179 variability is observed to be very large with ozone anomalies in excess of 35 μ g/m³ in 2003 180 relative to 2014. For 95th percentile ozone, the sharp increase by up to 20 μ g/m³ in the year 2003 181 occurred during a strong European heat wave (Section 4.2). The analysis of individual year 182 observations here shows that the increasing trend in the 5^{th} percentile ozone is a robust feature 183 184 with most of the recent years showing stronger noontime build up in ozone as compared to the 1990s. During the study period the variability in noontime ozone anomalies is however lower 185 $(\sim 10 \ \mu g/m^3)$ in the 5th percentile ozone compared to the 95th percentile ozone. 186

187 Consistently with the results obtained for hourly ozone, when the observational data is reduced to 188 diurnal values, a growth rate of $0.22 \pm 0.15 \ \mu g/m^3/y$ is calculated for the 5th percentile ozone, 189 while a stronger decline rate of $-0.57 \pm 0.34 \ \mu g/m^3/y$ is estimated for the 95th percentile ozone 190 (see Table 2). Hereafter we will mainly focus on trends in the daytime mean, nighttime mean, 5th 191 percentile and 95th percentile ozone concentrations.

The observed long-term reduction in 95th percentile ozone concentrations over Europe concurs 192 with the reduction in anthropogenic emissions of ozone precursors. Anthropogenic emissions of 193 NO_x and CO over Europe declined by 35% and 58%, respectively, as calculated from the 194 MACCity inventory. Slower rates of ozone reduction during nighttime are suggested to be 195 combined effects of reduced titration due to lower NO_x emissions, and an increase in the global 196 background ozone concentrations during this period, probably due to growing precursor 197 198 emissions in China since 1995, which has been predicted by Lelieveld and Dentener (2000) based on atmospheric chemistry - transport modeling, and corroborated by satellite observations 199 (Richter et al., 2005; Krotov et al., 2016). The effect of anthropogenic emissions is discussed in 200 201 more detail in the Section 4.1.

Fig. 3 further shows ozone trends for each month of the year. The slight growth rates in the 5th 202 percentile ozone are approximately equally distributed at the level of $0.1 \pm 0.12 \text{ }\mu\text{g/m}^3/\text{v}$. 203 Conversely, the monthly trends for the 95th percentile ozone are negative with a most rapid 204 decrease rate of -1.67 $\pm 0.4 \text{ }\mu\text{g/m}^3/\text{y}$ in August. For the 50th percentiles (mean) the seasonal cycle 205 of ozone trends decline unevenly from January to August, then pick up in the following months. 206 It leads to the fastest ozone growth in December when the ozone production is minor due to the 207 relatively lowest solar UV fluxes and temperatures, and the maximum ozone decline in August, 208 which is the photochemically most active month in Europe. In December, the 50^{th} (mean) 209 percentile ozone increases at a rate of 0.41 \pm 0.21 µg/m³/y (0.32 \pm 0.09 µg/m³/y), while a decline 210 rate of -0.40 $\pm 0.24 \ \mu g/m^3/y$ (-0.51 $\pm 0.13 \ \mu g/m^3/y$) is calculated in August. 211





Table 3 shows the trends in seasonal ozone concentrations over Europe analyzed separately for 212 day- and nighttime. The ozone concentrations show pronounced differences in trends over the 213 different seasons. The mean surface ozone in summer, averaged over the selected 93 sites, 214 declines at rates of -0.32 \pm 0.24 µg/m³/y and -0.20 \pm 0.27 µg/m³/y during day- and nighttime, 215 respectively. It is mainly related to the rapid decline in the highest levels (95th percentile) of 216 ozone with rates of -1.10 \pm 0.61 µg/m³/y (daytime) and -0.71 \pm 0.52 µg/m³/y (nighttime). 217 Although the 95th percentile ozone in spring declines almost as fast as during summer, the 218 decrease in spring for the 95th percentile ozone is compensated by the growth in 5th percentile 219 220 ozone, leading to much lower decrease rates in spring compared to summer for the mean ozone concentrations. Finally, in winter ozone grows at a rate of ~0.10 µg/m³/y. This increase occurs 221 mostly in the lower level (5th percentile) ozone concentrations, with growth rates of 0.25 ± 0.15 222 $\mu g/m^3/y$ (daytime) and 0.14 $\pm 0.22 \ \mu g/m^3/y$ (nighttime). 223

For the trends in annual mean ozone mixing ratios, a decline in the 95th percentile ozone (daytime: $-0.81 \pm 0.46 \ \mu g/m^3/y$; nighttime: $-0.57 \pm 0.36 \ \mu g/m^3/y$) is observed while an increase in the 5th percentile ozone (0.22 ± 0.17 and $0.16 \pm 0.17 \ \mu g/m^3/y$ for day- and nighttime, respectively, is calculated, resulting in statistically not-significant decreasing trends (daytime: -0.09 ± 0.24 ; nighttime: $-0.05 \pm 0.23 \ \mu g/m^3/y$) (Table 3).

Fig. 4 further shows the ozone trends distribution site-by-site over the 93 selected stations for 229 daytime mean, 5th and 95th percentile ozone during the four seasons. The 95th percentile ozone 230 trend shows a decline at most of the selected sites, although ozone increases are also visible at 231 several sites, especially in fall-to-winter. The annual ozone trend averaged over all sites during 232 daytime (-0.61 $\mu g/m^3/y$) is nearly twice that during nighttime (-0.34 $\mu g/m^3/y$, Fig. S2). For the 5th 233 percentile ozone, the annual means have grown over the western and central European sites, in 234 contrast with declines in ozone at other locations over the northern and southern Europe. 235 Averaged across all sites, the 5th percentile ozone has slightly grown during day- as well as 236 nighttime. The regional trend contrast is most significant in summertime with an average 237 decrease rate of -0.03 μ g/m³/y. The ozone trends spatial distribution in the daytime (Fig. 4) much 238 resembles that of the ozone trends in nighttime (Fig. S2) for the mean, 5th percentile as well as 239 95th percentile ozone. 240

241 **3.2 Ozone exceedance trends**

Based on the European directive for ozone concentrations limits, we calculate the number of exceedances for the information threshold and long-term objective (Fig. 5). Averaged over the selected 93 sites, the exceedances of the information threshold as well as the long-term objective have declined at rates of -3.2% and -2.5% per year relative to 1995. The decrease accelerated after the year 2003, during which a European heat wave raised summer temperatures by 20 to 30% (in degrees Celsius) compared to the seasonal average over a large part of the continent, extending from northern Spain to the Czech Republic and from Germany to Italy. The variations





in the exceedances are inter-annually consistent with the changes in the annual 95th percentile ozone, with a significant correlation coefficient of 0.93 for information threshold exceedances and 0.90 for long-term objective exceedances.

252 **3.3 Ozone trends from EMAC simulation**

The same analysis performed on the observations has been carried out on the EMAC model 253 results, i.e., for the same period covered by the observations. To ensure spatiotemporal 254 255 consistency with the EMEP data, modeled ozone concentrations are sampled at the times and locations of the measurements. Fig. 6 compares the time series of modeled and observed monthly 256 mean ozone over Europe. Although the model overestimates the measurements with a mean bias 257 of 4.3 μ g/m³ over the 1995–2014 period, the simulation results are highly correlated with 258 observed ozone, with a significant correlation coefficient of 0.91. The high bias may be explained 259 260 by the coarse grid resolution of 2.8 degrees that was applied, leading to the artificial dispersion of localized NO_x emissions, which optimizes NO_x concentrations over Europe with respect to 261 chemical O_3 formation, also noticed by Joeckel et al (2016). Such overestimation of the observed 262 263 ozone due to coarse model horizontal resolution has been reported by Lin et al. (2008) and Yan et al. (2014, 2016). The overestimation after 2010 becomes more evident (mean bias 5.4 µg/m³), 264 mostly due to the emissions used in the model version used, being prescribed up to the year 2005 265 and predicted in the subsequent period. The modeled ozone biases are slightly higher (mean bias: 266 5.2 μ g/m³ and 6.7 μ g/m³ for 1995–2014 and 2010–2014, respectively) compared to the observed 267 de-seasonalized time series. Nevertheless, EMAC model can reproduce the observed inter-annual 268 and seasonal variabilities of ozone, with statistically significant correlation coefficients at most 269 observation sites. For the diurnal, daytime as well as nighttime mean ozone averaged across the 270 271 93 sites, the model-observation correlation is 0.84-0.92 (0.62-0.70 for de-seasonalized time series). 272

The EMAC modeled ozone trends per hour are shown in Fig. 6. The agreement with the 273 observationally estimated trends is good, although the model tends to overestimate the trends by 274 0.12 $\mu g/m^3/y$, 0.23 $\mu g/m^3/y$, 0.08 $\mu g/m^3/y$, and 0.36 $\mu g/m^3/y$ for the mean, 5th, 50th and 95th 275 percentile ozone, respectively. The measured diurnal cycle of the ozone trends (Fig. 2) is well 276 captured by the EMAC model for the 5th and 95th percentile ozone concentrations. Consistently, 277 the modeled temporal evolutions (Fig. S3) of annual European 5th percentile ozone anomalies are 278 larger compared to the observations (~15 μ g/m³ versus ~10 μ g/m³ enhancements during 279 photochemical buildup of ozone at midday hours during 1990-2014), while being smaller for the 280 95th percentile ($\sim 21 \ \mu g/m^3$ versus $\sim 30 \ \mu g/m^3$). Further, the EMAC model reproduces the jump in 281 high level ozone concentrations during the year 2003 that was affected by a major heat wave. 282

For the diurnal mean values, averaged over Europe, the model produces higher growth rates for the 5th percentile ozone and weaker decrease rates for the 95th percentile ozone compared to the observed trends (Table 2). For the 50th percentile and mean ozone trends averaged over Europe,





the model shows statistically insignificant changes, similar to the observed trends (Table 2). Fig. S4 further shows the spatial distribution of the simulated diurnal ozone trends. It corroborates that central Europe experiences the highest growth rate for the averaged (also 50th percentile) and 5th percentile ozone concentrations, and the strongest reduction for the 95th percentile ozone during all seasons.

For the trends per month, the EMAC model reproduces the observed variability with statistically 291 significant correlation coefficients of 0.88–0.90 for the mean, 50th and 95th percentile ozone 292 trends (Fig. 3 and Fig. S5). Seasonally, for the 95th percentile ozone the modeled ozone trends are 293 much weaker than from measurements in all seasons except the autumn (Table 3). The decreased 294 higher level ozone is probably driven by the anthropogenic ozone precursor emission decline 295 over these years, which has been studied in previous work of ozone change drivers and 296 corroborated in Sect. 3.4 with a sensitivity simulation. For the 5th percentile ozone, especially for 297 the daytime period, the increasing trends are enhanced in the model results during all seasons 298 (Table 3). 299

300 4. Anthropogenic emissions and climate variability

301 4.1 Effects of anthropogenic emissions

A sensitivity simulation is conducted with constant global anthropogenic emissions to test the 302 sensitivity of observed European ozone to inter-annual variability in climate, by removing the 303 effects of anthropogenic emission changes. Consequently, the decline in European emissions (Fig. 304 S6) is removed from the EMAC model. With constant emissions, the modeled ozone shows a 305 slight increase at the midday hours for the 95th percentile and a slight decrease for the 5th 306 percentile, in contrast to the trends calculated from the control simulation. In the sensitivity 307 simulations no significant trend (less than 0.1 $\mu g/m^3/y$) for any hour of the day is found, and also 308 no contrast in ozone trends between the 5th and 95th percentiles (Fig. 7), which was well 309 reproduced by the control simulation. Therefore, it appears that both the decreases in 95th 310 percentile ozone and the enhancements in 5th percentile ozone are associated with the rapid 311 decline in the precursor gases anthropogenic emissions over Europe, notably of NO_X, prescribed 312 by the MACCity inventory (Fig. S6). These results reflect the effectiveness in controlling high-313 level ozone, but being unsuccessful in controlling the lower level ozone. Evidently, the 35% 314 reduction in NO_x emissions in Europe was not sufficient to achieve substantial reductions in 315 ozone, especially of background levels, which are affected by growing emissions in Asia that are 316 transported hemispherically (Lelieveld and Dentener, 2000). 317

Averaging over the selected 93 sites, we calculate the number of exceedances for the information threshold both in the control and the sensitivity simulation (Fig. 8). In the control simulation, the exceedances of the information threshold have declined at rates of -2.5% per year relative to 1995,





321 slightly smaller than the observed decrease rate of -3.2%. The variations in exceedances are inter-

annually consistent with the observations, with a significant correlation coefficient of 0.61. However, in the sensitivity simulation, the decline rate (-0.6%) in the exceedances is much

smaller than the rates in the control simulation and in the observations.

By fixing the anthropogenic emissions, ozone trends in each month for the 95th percentile ozone 325 show no obvious decline but rather a slight enhancement with growth rates of -0.23 - 0.50326 $\mu g/m^3/y$. For the 5th percentile ozone and compared to the control simulation, there is no increase 327 but a slight decrease at a rate of $-0.51 - 0.15 \ \mu g/m^3/y$ in months of the year (Fig. S7). For the 328 trends in annual mean ozone mixing ratios simulated in the sensitivity simulation, an 329 enhancement in the 95th percentile ozone (daytime: $0.16 \pm 0.18 \ \mu g/m^3/y$; nighttime: 0.10 ± 0.15 330 $\mu g/m^3/y$) is calculated while a decline in the 5th percentile ozone (-0.11 ± 0.14 and -0.07 ± 0.12 331 $\mu g/m^3/v$ for daytime and nighttime, respectively) is estimated, contrasting to but smaller in the 332 333 absolute value than the trends in the control simulation. This contrast has been also shown in the trends for individual hour of the day between control and sensitivity simulations (Fig. 7). These 334 results show that the effects of decline in anthropogenic emissions on European ozone change are 335 336 somewhat offset by the impacts of climate variability. This compensation effect is not only for 337 the high level ozone concentrations, which has been reported by previous studies (Lin et al., 2017), but also for the low level ozone concentrations. 338

339 **4.2 Effects of heat waves**

As discussed in number of studies (e.g., Filleul et al, 2006, Vautard et al, 2005, Garcia-Herrera et 340 a 2010, Vieno et al 2010), the 2003 heat waves caused favorable meteorology for ozone buildup, 341 342 leading to very high ozone concentrations during the summer period (from July to August). Fig. 9 shows the distribution of the difference in the exceedances between 2003 and averaged over 343 1995-2002 for the information threshold as well as the long-term objective over individual site. 344 Except for some northern sites, the exceedances in 2003 are much more frequent than the average 345 from 1995 to 2002 over most of the observational sites, especially over central Europe. This 346 exceedance anomaly distribution in 2003 relative to the period of 1995-2002 coincides with the 347 2-meter temperature anomaly distribution, with a statistically significant correlation up to 0.64 348 349 (Fig. S8).

The exceedance anomaly of information threshold and long-term objective during the year 2003 350 with respect to the 1995-2002 period follows the anomaly in ozone concentrations, in turn 351 consistent with the temperature anomaly. Fig. S9 shows the correlations between the monthly 352 mean 2-meter temperature and the monthly mean, 5th and 95th percentile ozone for diurnal, 353 daytime and nighttime concentrations. It corroborates the high correlations over central Europe 354 with statistically significant values up to ~ 0.82 . Ozone concentrations over the northern sites are 355 negative and insignificantly correlated with temperature. This may be related to the influence of 356 the Northern Atlantic Oscillation, which had an opposite impact on temperature over northern 357





compared to central Europe (Fig. S10). These results underscore that the large-scale climate
 variability affects the inter-annual variability of European ozone.

In the simulation with constant emissions, however, the modeled ozone fluctuation of annual 360 European ozone anomalies for individual hours is comparable in magnitude with the results in the 361 control simulation (Fig. S7). In both simulations, the fluctuation dominates around midday for 5th 362 (~15 μ g/m³ in the base simulation versus ~13 μ g/m³ in sensitivity simulation) and 95th (~21 363 $\mu g/m^3$ versus ~20 $\mu g/m^3$) percentile ozone (Fig. S7 and Fig. S3). In addition, the variations in the 364 365 exceedances of the information threshold are inter-annually consistent with the observations and the control simulation, with significant correlation coefficients of 0.54 and 0.56, respectively, 366 comparable to the correlations between observations and control simulation (Fig. 8). Further 367 correlations between the European averaged monthly mean 2-meter temperature and the modeled 368 monthly mean (50th), 5th and 95th percentile ozone in the sensitivity simulation are statistically 369 significant with correlation coefficients of 0.69-0.78 for diurnal, day- and nighttime 370 concentrations, consistent with the correlations (0.70-0.81) between 2-meter temperature and 371 simulated European ozone in the control simulation. These results clearly show that the ozone 372 373 variations are regulated by climate variations.

5. Conclusions and outlook

Based on EMEP observed ozone in the period 1995-2014, we analyzed the annual and seasonal 375 trends of the mean, the 5th, 50th and 95th percentile of the ozone concentrations at different 376 temporal distributions, i.e., hourly, diurnal, day- and nighttime. Results show that although 377 reductions in anthropogenic emissions have lowered the peak ozone concentrations, especially 378 during daytime in the period 1995-2014, the lower level ozone concentrations have increased 379 continually since 1995 over the 93 sites. This leads to insignificant trends in the 50th percentile 380 and mean ozone. Both the 5th and 95th percentile ozone trends follow a diel cycle with largest 381 trends during periods of strong photochemical activity. These contrasting ozone trends per hour 382 during the day and at different concentration levels are well reproduced by the EMAC chemistry-383 384 climate model, although the model slightly overestimates observed ozone at the surface. Furthermore, the numbers of exceedances of the information threshold and long-term objective 385 have continuously declined during the 20-year period considered, and the decrease has 386 387 accelerated since the year 2003.

Sensitivity simulations with constant emissions in the EMAC model, and correlation analysis 388 between modeled ozone and the ERA-Interim 2-meter temperature help distinguish effects of 389 390 climate and anthropogenic emissions on ozone variations and trends. Climate variability generally regulates the interannual variations of European surface ozone, while the changes in 391 anthropogenic emissions predominantly contribute to ozone trends. However, it appears that the 392 393 negative ozone trend due to European emission controls has been counteracted by a climate related tendency as well as hemispheric dispersion of pollutants from other regions, notably 394 China.. 395





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Fig. 1. Site distribution for the EMEP datasets (1990-2014) as well as the selected 93 sites (1995-2014). The overlaid

- Fig. 1. Site distribution for the EMEP datasets (1990-2014) as well as the selected 93 sites (1995-2014). The overlaid map shows the surface elevation (m) from a 2 min Gridded Global Relief Data (ETOPO2v2) available at NGDC
- map shows the surface elevation (m) from a 2 min Gridded Global Relief Data (ETOPO2v2) availa
 Marine Trackline Geophysical database (http://www.ngdc.noaa.gov/mgg/global/etopo2.html).
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Fig. 2. Trend in the observed surface ozone averaged over Europe, calculated for the selected 93 sites. The black line shows the 1995–2014 linear trends in the deseasonalized European monthly ozone anomalies for each hour of the day (local standard time), the red, purple and blue lines depict the observed trend for 5th, 50th and 95th percentile ozone, respectively, and the dashed bars indicate their standard deviations.









542 Fig. 3. Monthly trend in the observed surface ozone averaged over Europe for the selected 93 sites. The black line 543 shows the 1995-2014 linear trends in the European mean ozone for each month of the year, the red, purple and blue 544 lines depict the observed trend for 5th, 50th and 95th percentile ozone, respectively, and the dashed bars indicate their 545 standard deviations. The left axis is for the trends of mean, 5th, and 50th percentile ozone, while the right axis for the 546 95th percentile ozone.

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Fig. 4. Spatial distribution of measured daytime ozone trends in $\mu g/m^3$ across the selected 93 sites for average, 5th, 50th and 95th percentile ozone in annual mean and four seasons. Also shown in each panel are the average trends over all sites.







Year
 Fig. 5. Annual exceedances of the information threshold (for blue bars, hours should be multiplied by 100, 1-hourly

557 averages: 180 μ g/m³) as well as the long-term objective (red bars, maximum diurnal 8-hourly mean: 120 μ g/m³),

558 compared with the annual 95th percentile ozone concentrations (black line). Red dotted line shows the target value

(long-term objective that should not be exceeded more than 25 days per year, averaged over 3 years).







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Fig. 6. EMAC modeled ozone in µg/m³ over Europe during 1995-2014. Time series of measured (black) and modeled (red) monthly mean ozone over the 93 selected sites (top). Trend in the modeled surface ozone averaged over the selected 93 sites for all hours of the day (local time, bottom). The black line shows the 1995-2014 linear trends in the European mean ozone, the red, purple, and blue lines are the modeled trends for 5th, 50th and 95th percentile ozone, respectively. The dashed bars indicate their standard deviations.







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Fig. 7. Modeled trend in the surface ozone averaged over the selected 93 sites for all hours of the day (local time). The solid lines (left legends) show the 1995-2014 linear trends in the control simulation for 95th (top) and 5th percentile (bottom) ozone, respectively. The dashed lines (right legends) represent the modeled trends by the constant emission simulation. The bars indicate their deviations.





575



577 Fig. 8. Annual observed (top) and modeled (middle: control simulation; bottom: constant emission simulation) 578 exceedances of the information threshold (1-hourly averages: 180 μ g/m³). The hours along the y-axis should be 579 multiplied by 100.

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Fig. 9. Spatial distribution of the exceedance anomalies in 2003, relevant to the averages over 1995-2002 and for the information threshold as well as the long-term objective, in comparison with the 2-meter temperature anomalies in each of the sites.





Year	Number of sites	Missing data				
		Whole day	Daytime	Nighttime		
1995	113	32.6%	30.6%	34.6%		
1996	115	28.8%	26.7%	30.9%		
1997	121	23.9%	21.6%	26.2%		
1998	120	18.5%	16.0%	21.0%		
1999	127	10.4%	7.9%	12.8%		
2000	132	9.8%	7.2%	12.3%		
2001	134	11.9%	9.4%	14.4%		
2002	136	9.3%	6.8%	11.8%		
2003	137	12.1%	9.8%	14.4%		
2004	135	10.9%	8.5%	13.3%		
2005	132	10.5%	8.1%	12.9%		
2006	130	10.6%	8.1%	13.1%		
2007	132	9.5%	7.0%	12.0%		
2008	136	10.8%	8.2%	13.4%		
2009	134	10.6%	7.8%	13.3%		
2010	136	15.0%	12.6%	17.5%		
2011	135	13.8%	11.4%	16.2%		
2012	136	14.1%	11.8%	16.4%		
2013	136	19.9%	17.8%	22.0%		
2014	137	21.0%	19.1%	23.0%		

Table 1. Percentage of missing hourly data in each year.





- 590 Table 2. Modeled and observed ozone trends and their standard deviations based on diurnal
- average ozone concentrations. The mean, 5^{th} , 50^{th} , and 95^{th} percentile represent the monthly statistics of the diurnal averages. The model has been sampled in the same location of the EMEP
- 593 stations.

	5 th percentile	50 th percentile	Mean	95 th percentile
EMEP ($\mu g/m^3/y$)	0.22 ± 0.15	-0.05±0.23	-0.07±0.21	-0.57±0.34
EMAC (µg/m ³ /y)	0.42 ±0.14	0.01 ± 0.10	0.06 ±0.09	- 0 .23±0.10





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- 597 Table 3. Modeled and observed linear trends and their standard deviations of the 1995–2014
- 598 European mean annual and seasonal averaged daytime and nighttime mean as well as their 5th,

 $599 \quad 50^{\text{th}} \text{ and } 95^{\text{th}} \text{ percentile ozone concentrations.}$

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	Seasons	Mean		5 th percentile		50 th percentile		95 th percentile	
		EMEP	EMAC	EMEP	EMAC	EMMP	EMAC	EMÉP	EMAC
Daytime	Annual	$-0.09\pm$	$0.00\pm$	$0.22\pm$	$0.45\pm$	$-0.06\pm$	-0.01 \pm	-0.81 \pm	$-0.48\pm$
$(\mu g/m^3/y)$		0.24	0.06	0.17	0.14	0.24	0.06	0.46	0.15
	MAM	$-0.09\pm$	-0.05 \pm	$0.13\pm$	$0.52\pm$	-0.02 \pm	-0.02 \pm	-0.93 \pm	-0.49±
		0.27	0.08	0.24	0.17	0.27	0.08	0.53	0.16
	JJA	-0.32 \pm	-0.10±	-0.03 \pm	$0.41\pm$	-0.26±	$-0.09\pm$	-1.10±	$-0.54\pm$
		0.24	0.07	0.26	0.20	0.24	0.13	0.61	0.16
	SON	-0.03 \pm	-0.04 \pm	$0.09\pm$	$0.36\pm$	$-0.04\pm$	-0.02 \pm	-0.24 \pm	$-0.44\pm$
		0.19	0.05	0.14	0.12	0.20	0.05	0.25	0.23
	DJF	$0.10\pm$	$0.18\pm$	$0.25\pm$	$0.39\pm$	$0.05\pm$	$0.15\pm$	$-0.28\pm$	$-0.08\pm$
		0.25	0.14	0.15	0.22	0.27	0.20	0.31	0.05
Nighttime	Annual	-0.05 \pm	$0.12\pm$	$0.16\pm$	$0.38\pm$	-0.05 \pm	$0.07\pm$	$-0.57\pm$	-0.21±
$(\mu g/m^3/y)$		0.23	0.11	0.17	0.19	0.24	0.12	0.36	0.10
	MAM	$-0.06\pm$	$0.08\pm$	$0.18\pm$	$0.23\pm$	-0.00 \pm	$0.04\pm$	$-0.64\pm$	$-0.20\pm$
		0.29	0.10	0.23	0.23	0.29	0.08	0.43	0.12
	JJA	$-0.20\pm$	$0.06\pm$	$0.07\pm$	$0.36\pm$	-0.15±	$0.04\pm$	$-0.71\pm$	-0.36±
		0.27	0.14	0.24	0.22	0.28	0.14	0.52	0.21
	SON	-0.03 \pm	$0.06\pm$	$0.05\pm$	$0.19\pm$	-0.05 \pm	$0.04\pm$	$-0.21\pm$	-0.23 \pm
		0.21	0.10	0.12	0.16	0.23	0.11	0.24	0.19
	DJF	$0.09\pm$	$0.24\pm$	$0.14\pm$	$0.43\pm$	$0.06\pm$	$0.20\pm$	-0.24±	$-0.05\pm$
		0.24	0.18	0.22	0.27	0.25	0.25	0.29	0.06