



1 **Cryosphere: a kingdom of anomalies and diversity**

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11 **Abstract.** The cryosphere of the Earth overlaps with the atmosphere, hydrosphere and lithosphere over vast areas
12 with temperatures below zero C and pronounced H₂O phase changes. In spite of its strong variability in space and
13 time, the cryosphere plays the role of a global thermostat keeping the thermal regime on the Earth within rather
14 narrow limits affording continuation of the conditions needed for the maintenance of life. Objects and processes
15 related to cryosphere are very diverse due to the following basic reasons: anomalous thermodynamic and
16 electromagnetic properties of H₂O, intermediate intensity of hydrogen bonds, and very wide spread of cryogenic
17 systems all over the Earth. These features yet attract insufficient attention of research communities. Cryology is
18 usually understood as a descriptive discipline within physical geography basically limited to glaciology and
19 permafrost research. We emphasize its broad interdisciplinary landscape involving physical, chemical and
20 biological phenomena related to the H₂O phase transitions and various forms of ice. This paper aims to attract



21 attention of readers to crucial importance of cryogenic anomalies which make the Earth atmosphere and the entire
22 Earth system very specific, if not unique, objects in the universe.

23 **1. Introduction**

24 Nowadays the Earth system is facing the so-called “Grand Challenges”. The rapidly growing population needs
25 fresh air and water, more food and more energy. Herewith, the humankind suffers from the climate change,
26 deterioration of the air, water and soils, deforestation, acidification of ocean waters and the biodiversity losses.
27 The Grand Challenges are threatening the security of the upcoming societies. As stated in the Earth System
28 Manifesto, the humankind has only a 40-year window to avoid the Earth system collapse
29 (https://www.atm.helsinki.fi/peex/images/manifesti_peex_ru_hub2.pdf).

30 Especially strong environmental changes are observed and, moreover, are expected to go on during the next
31 decades in the North Eurasia and Arctic Ocean (IPCC, 2014). The rate of the changes in the Arctic Region is
32 higher than elsewhere in the world (Smith et al., 2015; IPCC, 2013). Furthermore, the threats from climate
33 change and deterioration of the environment are redoubled by the restricted natural resources, unrestrained
34 migration tendencies, and uncertainties in political and socio-economic developments (Smith, 2010).

35 The totality of environmental problems in their relation to human activities in North Eurasia (the area extended
36 north of 45° N, including the Arctic Ocean and between the Atlantic and Pacific Oceans in latitude) is addressed
37 in the recently launched international program “Pan-Eurasian Experiment” (PEEX) aimed at responding to the
38 Grand Challenges faced by the European countries, Russia and China (Lappalainen et al., 2014; Kulmala et al.,
39 2015). The PEEX Science Plan (Lappalainen et al., 2016) focuses on the Polar and Arctic regions and,
40 specifically, on the cryosphere in the context of modern challenges of the North-Eurasian environment.

41 Historically, the cryosphere research has been developing in frames of physical geography, with the major focus
42 on glaciology and permafrost research. The present paper attracts attention of the broad geoscience community of



43 physicists, chemists, biologists and, prospectively, astrophysicists to topical problems related to ice, snow and
44 cold in a wider interdisciplinary context, beyond the present conventional understanding (see above).

45 **2. Cryosphere and cryogenic anomalies**

46 Cold regions and natural phenomena related to cold are often described with terms enclosing the prefix *cryo* that
47 means *cold*. The totality of phenomena on the Earth related to or linked with snow, ice or ice-like products (such
48 as gas hydrates) comprises the Earth's cryosphere – a complex synergic system with self-organized constituents
49 which maintains the heat and mass exchange with the environment and spreads over the atmosphere, the
50 hydrosphere, the lithosphere and the biosphere (Melnikov and Gennadinik, 2011).

51 The large extent of cryogenic objects, their variability in space and time, as well as the discontinuity of the
52 cryosphere, have inspired development of special management technologies. Such technologies, taking into
53 account the stabilizing and bio-protective effects of ice, rooted in specific properties of ice, have become
54 applicable to the areas outside the high-latitude and low temperature regions (Melnikov et al., 2010).

55 Recent investigations of the cryosphere have yielded discovery and scientific description of a number of unique
56 natural objects and phenomena called “cryogenic anomalies”. Below we consider examples of recently
57 investigated cryogenic anomalies, tentatively divided into six groups: ice structure, glaciation, ocean, atmospheric
58 ice and water, biota and cold, and rates of processes.

59 **3. Ice structure**

60 Being a compositionally simple material, ice possesses diverse anomalous properties. It can exist in seventeen
61 phase states, out of which eleven states are clearly expressed. This diversity shows up in physicochemical and
62 biological processes, and in the types of precipitation: one type for rain, eight types for snow, and two mixed
63 water-ice types (Eisenberg and Kauzman, 1975). There are paradoxically different properties brought together in



64 ice: it is at the same time elastic and plastic, crystalline and amorphous, semiconducting and dielectric, lighter
65 than water but harder than steel (Maeno, 1988). With its crystals built uniquely by hydrogen bonds, ice provides a
66 standard for estimating these bonds. The complexity of the ice structure and its non-equilibrium phase transitions
67 are sufficient for self-organizing synergetic behavior and formation of stable macroscopic objects, like the
68 classical snowflakes or drop clusters in atmospheric clouds (Shavlov et al., 2011).

69 **4. Glaciation**

70 Glaciation is the most impressive cryogenic phenomenon on the Earth. Currently ice covers more than 10% of the
71 Earth's surface, but the area of glaciation was much larger in some periods of the Earth's history. Traces of early
72 Proterozoic glaciations (Huron, about 2300 myr ago) are found on all modern continents, which is evidence of
73 global glaciation, even taking into account the drift of continents. About 300 myr ago a large part of Gondwana
74 was covered by ice (The Winters, 1979).

75 The exact cause of glaciations remains debatable but their cyclic nature is beyond doubt. The complex pattern of
76 glacial cycles cannot be explained uniquely by orbital forcing of eccentricity, obliquity, and precession that
77 change in Milankovitch cycles (Milanković, M., 1939), but correlates also with global geological events.

78 Glaciation obviously reduces the area of rocks exposed to wind erosion and causes progressive CO₂ increase by
79 more active respiration of bacteria, which recycle dead organics at less active photosynthesis. This, in turn,
80 enhances the greenhouse effect of the atmosphere and thus increases the temperature. Due to this negative
81 feedback, oscillations in the cryosphere correlate with phases of global greenhouse-gas and temperature cycles.
82 There is also a positive feedback that controls ice waxing and waning (Le Hir et al., 2008). In particular, the
83 increasing albedo may have been responsible for the strongest glaciation about 800 myr ago, the so called
84 Snowball Earth (Kirschvink et al., 1997). The observed recent shrinking of areas covered by glaciers is often



85 attributed to anthropogenic impacts which became a real geological force in the 20th century (Vernadsky V.I.,
86 1965). The true extent of this effect will be clear in the near future.

87 The very existence of ice sheets changes the composition and physical properties of the lithosphere, the
88 hydrosphere, and the atmosphere. Moving ice transports great amounts of material and, at the same time,
89 preserves the entrained objects. During glaciations, crust becomes denser; volcanism increases; ice sheets
90 extended to the troposphere change the atmospheric composition and circulation; sea level falls and climate
91 changes. After the ice retreat the formerly frozen territories experience isostatic rebound and terrain smoothing,
92 changes in landscapes, such as formation of rivers and lakes, and changes in biota, such as species composition.

93 **5. Permafrost**

94 Formation of permafrost is the principal consequence of glaciation. Remnants of permafrost that formed during
95 the last glacial event currently cover large territories. The stability of permafrost is one of the greatest challenges,
96 first of all for Russia, where it occupies 65% of the territory, but also for the global climate-biosphere system,
97 because of strong links between the permafrost and boreal forests, tundra, peatlands, rivers and other water
98 reservoirs. Objects and phenomena associated with permafrost to a large extent determine the sustainability of the
99 Siberian climate-biosphere system (taiga) and the conditions of land use and life of the population. The ice
100 barriers dramatically change natural flows of water and material in mid latitudes, change hydrological regimes
101 and create traps for gases. The methane traps cause jet degassing of hydrocarbons (Melnikov et al., 1997) which,
102 in turn, can lead to cooling of rocks, formation of gas hydrates, accumulation of large amounts of methane, and to
103 the methane explosions. The latter cause significant changes in both underwater and land topography (see Figure
104 2).



105 On the land, such explosions may produce giant craters which subsequently transform into lakes (see Figure 3).
106 Tracking such processes over large areas in real time with proper resolution in space, as well as the early warning
107 of expected gas explosions, are important tasks envisaged in the PEEEX Science Plan (Lappalainen et al., 2016).
108 Permafrost leaves imprints in the gravity field, as in the case of the Hudson Bay, where the gravity field is lower
109 than elsewhere at this latitude (Tamisiea et al., 2007), possibly, as a consequence of the North American
110 glaciation. About 75,000 years ago the ice (3.7 km thick in the Hudson Bay area) covered most of the present
111 northern US and Canada, which has left a large depression. On the contrary, the gravity field over the 1 km thick
112 permafrost (such as in East Siberia) can reach 50 mGal. On the geological time scale, this field can influence the
113 homeostasis of living organisms and their speciation.

114 **6. Ocean**

115 Oceanic ice is a major control of the global climate as it screens heat and mass transfer in both vertical (across the
116 air-water interface) and horizontal directions. Warm currents, such as that of North Atlantic, move mainly along
117 the ocean surface due to thermal expansion of warm water. Glaciers on the way of warm currents can melt
118 rapidly because the ice not necessarily should reach the seafloor to stop them but can be only commensurate with
119 the surface water layer.

120 Another great challenge addressed in the PEEEX Science Plan is the ongoing change in the Arctic Ocean ice
121 cover, showing increasingly larger and longer-exposed ice-free areas (Hayes et al., 2014; Schaefer et al., 2014;
122 Döscher et al., 2014). This tendency is favorable for exploring and transportation of the Northern Siberian oil and
123 gas comprising about 25% of their total world resources (Yenikeyeff and Krysiak, 2007).



124 **7. Atmospheric ice and water**

125 Ice has very high heat capacity, specific heat of melting, and dielectric constant. The phase transition forms
126 shields of different types and scales from the gas hydrates stored in permafrost to the ice sheets that cover the
127 entire Antarctic continent and affect the global climate and evolution of life. In the global Earth system, the
128 cryosphere acts as the temperature stabilizer due to exceptional thermal inertia of ice and water occupying a great
129 part of the Earth's surface.

130 The Earth's atmosphere contains a large amount of water in the form of vapor, liquid and ice which influences
131 the planetary heat budget. Moreover, the atmosphere acts as a "thermal shield" aggregating the water molecules
132 into ice particles and thus holding back on the Earth the water, vital for the maintenance of life. The phase
133 transitions of atmospheric ice produce diverse sudden effects and various precipitation types: rain, ice, drizzle,
134 moist snow, snow, sleet, or hail. In the dry troposphere, the air cools down adiabatically at 0.98°C per 100 m.
135 Hence, condensation of originally nearly saturated water vapor starts already at a few hundred meters above the
136 surface. Along with sublimation, this leads to the heat release, which affects temperature profiles and other
137 properties of the atmosphere. Generally, condensation of pure water vapor starts at 400 - 00% supersaturation
138 (e.g. Pruppacher and Klett, 1997). However, condensation in the real atmosphere starts shortly after relative
139 humidity exceeds 100% due to the presence of solid or liquid aerosol particles acting as condensation nuclei, or
140 even below 100% in some cases (Kulmala et al., 1997). Similarly, ice nucleation is initiated upon nuclei like dust
141 particles, biological particles etc., including anthropogenic aerosol particles (e.g. Hoose and Möhler, 2010,
142 Atkinson et al. 2013). Thus, the air pollution facilitates formation of ice crystals and cloud droplets, and therefore
143 affects the air temperature and precipitation processes and its spatiotemporal variability (Rosenfeld et al. 2014).
144 The aerosol-cloud-climate interactions (see e.g. Kulmala et al., 2011), as well as air pollution-weather-climate
145 interactions (Ding et al. 2016, Petäjä et al., 2016) have been and are under rigorous investigations.



146 **8. Biota and cold**

147 A number of facts in the evolution biology would be unexplainable without cryosphere. Contrary to the common
148 thinking, the existence of life is consistent with negative temperatures. For example, photosynthesis is quite
149 possible at temperatures below 0°C (Kramer P.J., Kozłowski T.T., 1960). For billions of years, the cryosphere
150 assured the stability of environment required for development of life. The mean annual air temperatures on the
151 Earth were naturally maintained near the ice-water transition point. Ice, like water, possesses exceptional
152 thermostatic properties, and is thus a very strong stabilizer of the temperature regime, along with wide spread of
153 ice over the Earth's surface. The heat capacity of water (4.183 kJ/kg·K) is five times the mean heat capacity of
154 soil, and its volumetric heat capacity is 3300 times that of air. The high heat capacity of ice (2.06 kJ/kg·K) is
155 comparable with that of water, which makes the ice-snow-water system the Earth's largest storage of solar
156 energy. The very phase change point has additional (likewise abnormal) thermal stability: the specific heat of ice
157 melting is five times as high as in gold (332 against 66.2 kJ/kg) and is higher than, say, in mercury by a factor 28
158 (Melnikov and Gennadinik, 2011).

159 Permafrost contains organic matter in the form of biogenic gases, products of biodegradation, or as ancient viable
160 bacteria. In Russia, the evidence of bacteria in permafrost was obtained in the latest 19th century, together with
161 finding of mammoths in Northern Siberia and during studies of soils in the Russian Far East. Novel information
162 about bacterial communities has been obtained from studies of Antarctic permafrost. Cyanobacteria, with their
163 age about 500 kyr, were discovered in the station "Vostok" ice cores at 3600 m depth (Water sampling, 2003).
164 Many microorganisms are stable against freezing and not only survive but even grow at negative temperatures.
165 More than 10% of water in organic tissues remains unfrozen at temperatures below -20°C. Many basic processes
166 in microbial, soil, geophysical, and cosmological systems are associated with cooling and ice-water phase
167 changes. Ribosome activity is often more successful at low temperatures, which may be the evidence that life
168 could have originated in cold conditions (Vlassov et al., 2005).



169 Life processes can remain active at below 0 ° C in capillaries and on ice surface in permafrost. Micro-organisms
170 can exist in media with low dimensionality (2 or less), such as liquid films and capillaries filled with super-
171 cooled water. The porous structure of the environment ensures the transport of nutrients and life wastes
172 (Möhlmann, D.T.F., 2009).

173 Even comparatively minor heating to 50-60°C leads to denaturation of proteins and can stop operation of a living
174 system, while freezing even to the near-absolute zero temperatures does not change the configuration of
175 biomolecules, so that the life functions can resume after thawing. Enzymes that regulate metabolism in organisms
176 show similar behavior (Dethier and Villed, 2005).

177 It is still unclear how and why the metazoans originated about a billion years ago. The obviously required
178 sufficient amount of oxygen has been commonly explained by its gradual photosynthetic accumulation in the
179 atmosphere. However, origination of metazoans may have been due to the higher concentration of oxygen in cold
180 water. The origin of the skeletal organisms at the Proterozoic/Phanerozoic boundary was possibly favored by the
181 calcium carbonate release by organisms at the high oxygen availability. The oxygen enrichment in oceans at the
182 Vendian/Cambrian boundary may have resulted from glaciation, water cooling, hence, increasing gas solubility
183 and favorable water circulation patterns. Note that the first giant organisms have appeared precisely during this
184 cooling event. Furthermore, gigantism, associated with specific heat exchange, is known to be typical of polar
185 biota. Biotic and cooling events have been closely related all over the Phanerozoic. The major Jurassic and latest
186 Cretaceous extinctions occurred during the stages of cold climate (The Winters, 1979).

187 **9. Typical rates of processes**

188 Natural objects are characterized by their typical scales in time and space. The common sense suggests that
189 temporal and spatial scales change more or less similarly, which implies that typical change rates are not too
190 variable. However, this rule does not work for the cryosphere, which can be either accelerator or decelerator of



191 various processes (Melnikov et al., 2013). For instance, paleo-bacteria preserved under ice shields remain viable
192 for millions of years. Ice slow down biotic processes due to the following properties:

- 193 • low dielectric and magnetic susceptibility, which reduces electromagnetic fields;
- 194 • high stiffness, which resists mechanic impacts;
- 195 • low permeability, which decelerates mass transport;
- 196 • high heat capacity and anomalous specific heat of melting, which damp temperature variations.

197 Hence, interaction between an object and its environment in cryosphere is basically restricted to weak gravity
198 effects and very slow diffusion. The capability of ice to decelerate biological processes has practical uses:
199 freezing of organs or storage of seeds, not to mention the household refrigerators. Similar mechanisms protect the
200 biota of Antarctic paleo-lakes, such as Lake Vostok, hidden under kilometers-thick ice maintaining its sustainable
201 thermal regime (Thoma et al., 2008).

202 On the contrary, processes on the surface of ice or ice-bearing systems are often faster than usual due to
203 formation of electrolytes, chemical reactions, and accelerated transport. “Ice fingers of death” is an example of a
204 very rapid process with extremely high gradients of physical parameters in the vicinity of cryogenic objects. This
205 phenomenon (also known as brinicles, brine icicles or ice stalactites) develops in shallow water beneath the sea
206 ice, mainly in the southern high latitudes, in the form of slowly rotating jets of extremely cold and saline (and
207 therefore heavy) surface water sinking down to the ocean bottom (Martin, 1974). The point is that the formation
208 of sea ice at very low air temperatures is associated with the decomposition of calcium bicarbonate into CO₂ and
209 CaCO₃. Then CO₂ releases into the air, whereas slowly rotating jets of the cold dense surface water sinks down
210 to the seafloor and propagates downslope leaving frozen footprints with entrapped bottom-dwelling sea animals.

211 Cryogenic phenomena developing at various rates are presented in the log-log space/time coordinates in Fig. 4
212 showing extremely wide range of scales. Some objects, such as paleo-bacteria survived in ice over half a million



213 years or “ice fingers of death”, fall beyond the scope of any particular science. To emphasize their
214 multidisciplinary nature, these objects are marked by two colors indicating the relevant pairs of disciplines.

215 **10. Cryodiversity and its role in the Earth system**

216 The Earth’s cryosphere spreads from the oceanic depths reaching ~5 km to the boundary between the atmosphere
217 and outer space (called Karman line) at ~100 kilometers above the sea level. The cryospheres of the Earth and
218 other planets with their satellites actually comprise the cryosphere of the Sun System with its radius ~100 au,
219 where the density of solar wind (radiated plasma) is high. Thus, the terrestrial cryosphere is an object among a
220 diversity of planets and other large objects in the Solar System, which differ in their distances to the Sun; rotation
221 rates and orbiting paths with the respective diurnal and annual cycles; chemical compositions; and presence or
222 absence of the atmospheres. Some planets of the Solar System and their satellites store great amounts of ice,
223 much more than the Earth (Fig. 5), both in absolute values and relative to the mass of the respective celestial
224 bodies. Recent extraterrestrial missions have demonstrated impressive cryogenic processes and phenomena on
225 various planets and satellites (Komarov and Isaev, 2010).

226 **11. Cryogenic objects in the Solar system**

227 Historically, the Earth’s cryosphere was investigated in the scope of exploration of the high-latitude or mountain
228 regions characterized by cold climate, large stocks of ice and snow, and strong thermal effects of the H₂O phase
229 transitions. Those investigations were basically performed in the conceptual framework of geographical
230 disciplines, such as glaciology or permafrost science. The first priority belonged to description, quantification and
231 inventory of phenomena, with comparatively little attention to concrete features of the background physical,
232 chemical and biological processes. Such approach, although motivated by immediate practical purposes, limits
233 our knowledge of the cryosphere basically to its present state and geological past.



234 In view of the ongoing climate change and massive anthropogenic deterioration of the environment, it becomes
235 critical to be able to predict the near future of the Earth's cryosphere in its capacity as the major stabilizer of the
236 Earth climate and environment. To this end, a leap is needed from the currently dominating descriptive approach
237 to the process-based approach unifying investigation of the totality of cryogenic phenomena from molecular to
238 planetary scales (see Figure 4) in an interdisciplinary science, *cryology*, armed with all necessary methods of
239 geosciences, physics, chemistry and biology.

240 **12. Concluding remarks**

241 Ice has very high heat capacity, specific heat of melting, and dielectric constant. The phase transition forms
242 shields of different types and scales from the gas hydrates stored in permafrost to the ice sheets that cover the
243 entire Antarctic continent and affect the global climate and evolution of life. In the global Earth system, the
244 cryosphere acts as the temperature stabilizer due to exceptional thermal inertia of ice and water occupying a great
245 part of the Earth's surface.

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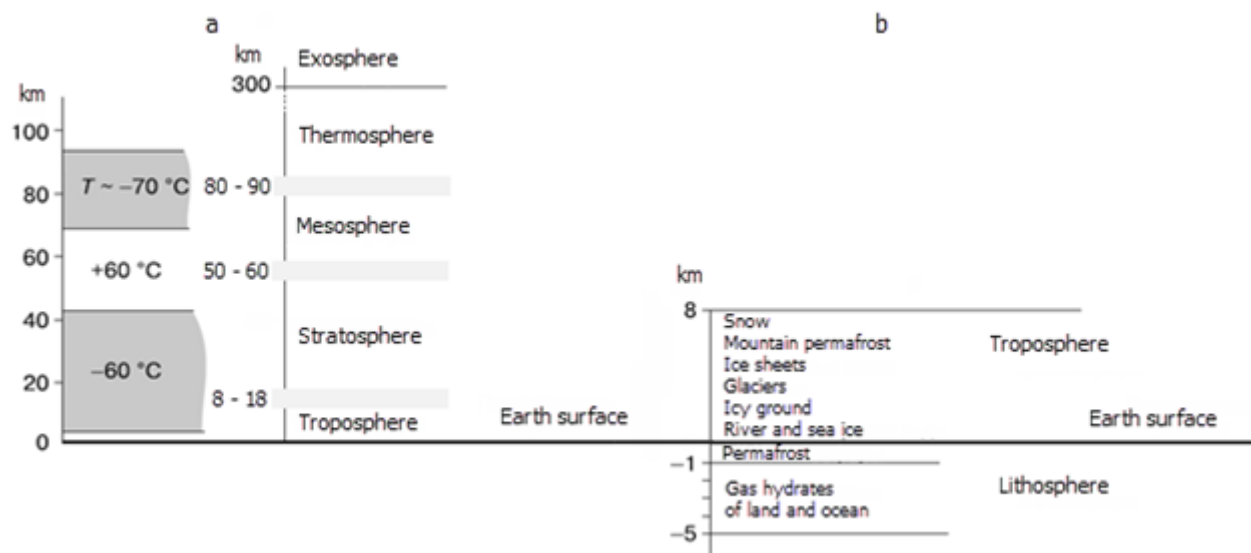
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428 Figure 1: The Earth's cryosphere: from (a) the uppermost atmosphere, to (b) 5 km below Earth's surface, where

429 ice exists as hydrates at positive temperatures.



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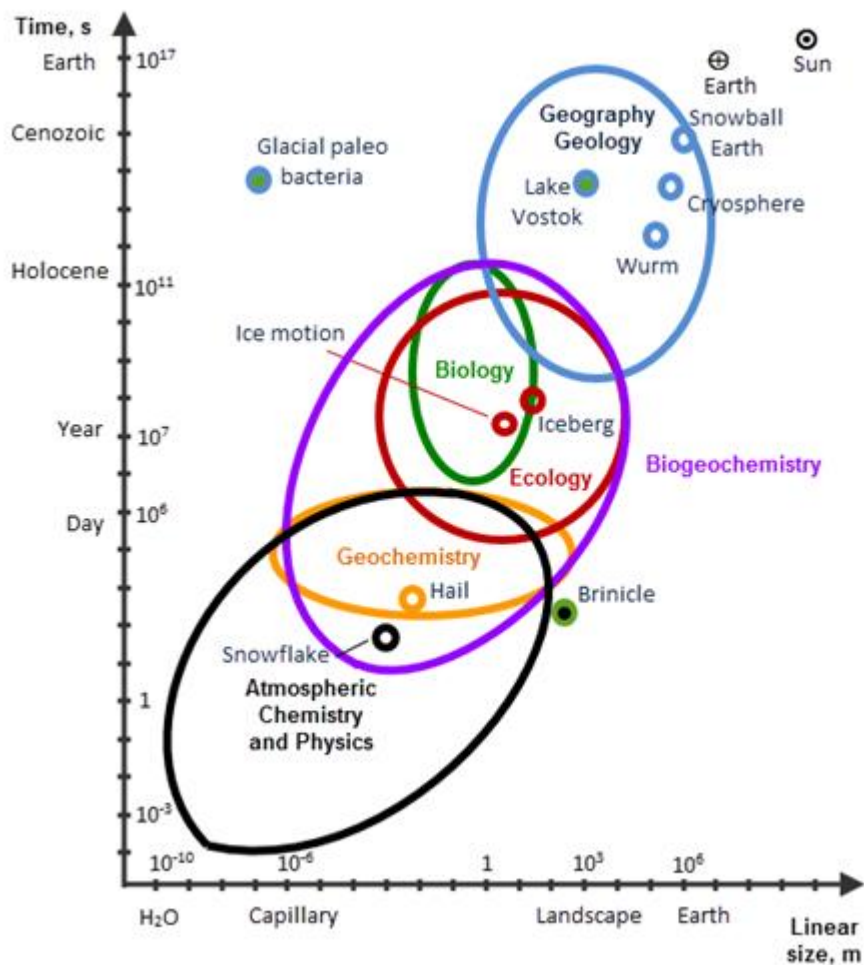
431 Figure 2: Photo of a lake on the Yamal Peninsula showing numerous craters from gas explosions on the lake

432 bottom (Bogoyavlensky, 2015).



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434 Figure 3: New crater produced by gas explosion on the Yamal Peninsula in 2014 (Leybman et al., 2015;
435 photograph by Vladimir Pushkarev borrowed from “Siberian Times”).



436

437 Figure 4: Log-log space-time diagram of cryogenic phenomena of different nature. Large ellipses or circles with
438 different colors identify various sciences. Small circles show various objects of the cryosphere. The Earth and the
439 Sun are shown to illustrate the range of scales.