

1 **Cryosphere: a kingdom of anomalies and diversity**

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12 **Abstract.** The cryosphere of the Earth overlaps with the atmosphere, hydrosphere and lithosphere over vast areas
13 with temperatures below zero C and pronounced H₂O phase changes. In spite of its strong variability in space and
14 time, the cryosphere plays the role of a global thermostat keeping the thermal regime on the Earth within rather
15 narrow limits affording continuation of the conditions needed for the maintenance of life. Objects and processes
16 related to cryosphere are very diverse due to the following basic reasons: anomalous thermodynamic and
17 electromagnetic properties of H₂O, intermediate intensity of hydrogen bonds, and very wide spread of cryogenic
18 systems all over the Earth. These features yet attract insufficient attention of research communities. Cryology is
19 usually understood as a descriptive discipline within physical geography basically limited to glaciology and
20 permafrost research. We emphasize its broad interdisciplinary landscape involving physical, chemical and
21 biological phenomena related to the H₂O phase transitions and various forms of ice. This paper aims to attract

22 attention of readers to crucial importance of cryogenic anomalies which make the Earth atmosphere and the entire
23 Earth system very specific, if not unique, objects in the universe.

24 **1. Introduction**

25 Nowadays the Earth system is facing the so-called “Grand Challenges”. The rapidly growing population needs
26 fresh air and water, more food and more energy. Herewith, the humankind suffers from the climate change,
27 deterioration of the air, water and soils, deforestation, acidification of ocean waters and the biodiversity losses. The
28 Grand Challenges are threatening the security of the upcoming societies. As stated in the Earth System Manifesto,
29 the humankind has only a 40-year window to avoid the Earth system collapse
30 (https://www.atm.helsinki.fi/peex/images/manifesti_peex_ru_hub2.pdf).

31 Especially strong environmental changes are observed and, moreover, are expected to go on during the next decades
32 in the North Eurasia and Arctic Ocean (IPCC, 2014). The rate of the changes in the Arctic Region is higher than
33 elsewhere in the world (Smith et al., 2015; IPCC, 2013). Furthermore, the threats from climate change and
34 deterioration of the environment are redoubled by the restricted natural resources, unrestrained migration
35 tendencies, and uncertainties in political and socio-economic developments (Smith, 2010).

36 The totality of environmental problems in their relation to human activities in North Eurasia (the area extended
37 north of 45° N, including the Arctic Ocean and between the Atlantic and Pacific Oceans in latitude) is addressed in
38 the recently launched international program “Pan-Eurasian Experiment” (PEEX) aimed at responding to the Grand
39 Challenges faced by the European countries, Russia and China (Lappalainen et al., 2014; Kulmala et al., 2015).
40 The PEEX Science Plan (Lappalainen et al., 2016) focuses on the Polar and Arctic regions and, specifically, on the
41 cryosphere in the context of modern challenges of the North-Eurasian environment.

42 Historically, the cryosphere research has been developing in frames of physical geography, with the major focus
43 on glaciology and permafrost research. The present paper attracts attention of the broad geoscience community of

44 physicists, chemists, biologists and, prospectively, astrophysicists to topical problems related to ice, snow and cold
45 in a wider interdisciplinary context, beyond the present conventional understanding (see above).

46 **2. Cryosphere and cryogenic anomalies**

47 Cold regions and natural phenomena related to cold are often described with terms enclosing the prefix *cryo* that
48 means *cold*. The totality of phenomena on the Earth related to or linked with snow, ice or ice-like products (such
49 as gas hydrates) comprises the Earth's cryosphere – a complex synergic system with self-organized constituents
50 which maintains the heat and mass exchange with the environment and spreads over the atmosphere, the
51 hydrosphere, the lithosphere and the biosphere (Melnikov and Gennadinik, 2011).

52 The large extent of cryogenic objects, their variability in space and time, as well as the discontinuity of the
53 cryosphere, have inspired development of special management technologies. Such technologies, taking into
54 account the stabilizing and bio-protective effects of ice, rooted in specific properties of ice, have become applicable
55 to the areas outside the high-latitude and low temperature regions (Melnikov et al., 2010).

56 Recent investigations of the cryosphere have yielded discovery and scientific description of a number of unique
57 natural objects and phenomena called “cryogenic anomalies”. Below we consider examples of recently investigated
58 cryogenic anomalies, tentatively divided into six groups: ice structure, glaciation, ocean, atmospheric ice and water,
59 biota and cold, and rates of processes.

60 **3. Ice structure**

61 Being a compositionally simple material, ice possesses diverse anomalous properties. It can exist in seventeen
62 phase states, out of which eleven states are clearly expressed. This diversity shows up in physicochemical and
63 biological processes, and in the types of precipitation: one type for rain, eight types for snow, and two mixed water-
64 ice types (Eisenberg and Kauzman, 1975). There are paradoxically different properties brought together in ice: it

65 is at the same time elastic and plastic, crystalline and amorphous, semiconducting and dielectric, lighter than water
66 but harder than steel (Maeno, 1988). With its crystals built uniquely by hydrogen bonds, ice provides a standard
67 for estimating these bonds. The complexity of the ice structure and its non-equilibrium phase transitions are
68 sufficient for self-organizing synergetic behavior and formation of stable macroscopic objects, like the classical
69 snowflakes or drop clusters in atmospheric clouds (Shavlov et al., 2011).

70 **4. Glaciation**

71 Glaciation is the most impressive cryogenic phenomenon on the Earth. Currently ice covers more than 10% of the
72 Earth's surface, but the area of glaciation was much larger in some periods of the Earth's history. Traces of early
73 Proterozoic glaciations (Huron, about 2300 myr ago) are found on all modern continents, which is evidence of
74 global glaciation, even taking into account the drift of continents. About 300 myr ago a large part of Gondwana
75 was covered by ice (The Winters, 1979).

76 The exact cause of glaciations remains debatable but their cyclic nature is beyond doubt. The complex pattern of
77 glacial cycles cannot be explained uniquely by orbital forcing of eccentricity, obliquity, and precession that change
78 in Milankovitch cycles (Milanković, M., 1939), but correlates also with global geological events.

79 Glaciation obviously reduces the area of rocks exposed to wind erosion and causes progressive CO₂ increase by
80 more active respiration of bacteria, which recycle dead organics at less active photosynthesis. This, in turn,
81 enhances the greenhouse effect of the atmosphere and thus increases the temperature. Due to this negative feedback,
82 oscillations in the cryosphere correlate with phases of global greenhouse-gas and temperature cycles. There is also
83 a positive feedback that controls ice waxing and waning (Le Hir et al., 2008). In particular, the increasing albedo
84 may have been responsible for the strongest glaciation about 800 myr ago, the so called Snowball Earth (Kirschvink
85 et al., 1997). The observed recent shrinking of areas covered by glaciers is often attributed to anthropogenic impacts

86 which became a real geological force in the 20th century (Vernadsky V.I., 1965). The true extent of this effect will
87 be clear in the near future.

88 The very existence of ice sheets changes the composition and physical properties of the lithosphere, the
89 hydrosphere, and the atmosphere. Moving ice transports great amounts of material and, at the same time, preserves
90 the entrained objects. During glaciations, crust becomes denser; volcanism increases; ice sheets extended to the
91 troposphere change the atmospheric composition and circulation; sea level falls and climate changes. After the ice
92 retreat the formerly frozen territories experience isostatic rebound and terrain smoothing, changes in landscapes,
93 such as formation of rivers and lakes, and changes in biota, such as species composition.

94 **5. Permafrost**

95 Formation of permafrost is the principal consequence of glaciation. Remnants of permafrost that formed during the
96 last glacial event currently cover large territories. The stability of permafrost is one of the greatest challenges, first
97 of all for Russia, where it occupies 65% of the territory, but also for the global climate-biosphere system, because
98 of strong links between the permafrost and boreal forests, tundra, peatlands, rivers and other water reservoirs.
99 Objects and phenomena associated with permafrost to a large extent determine the sustainability of the Siberian
100 climate-biosphere system (taiga) and the conditions of land use and life of the population. The ice barriers
101 dramatically change natural flows of water and material in mid latitudes, change hydrological regimes and create
102 traps for gases. The methane traps cause jet degassing of hydrocarbons (Melnikov et al., 1997) which, in turn, can
103 lead to cooling of rocks, formation of gas hydrates, accumulation of large amounts of methane, and to the methane
104 explosions. The latter cause significant changes in both underwater and land topography (see Figure 2).
105 On the land, such explosions may produce giant craters which subsequently transform into lakes (see Figure 3).
106 Tracking such processes over large areas in real time with proper resolution in space, as well as the early warning
107 of expected gas explosions, are important tasks envisaged in the PEEEX Science Plan (Lappalainen et al., 2016).

108 Permafrost leaves imprints in the gravity field, as in the case of the Hudson Bay, where the gravity field is lower
109 than elsewhere at this latitude (Tamisiea et al., 2007), possibly, as a consequence of the North American glaciation.
110 About 75,000 years ago the ice (3.7 km thick in the Hudson Bay area) covered most of the present northern US and
111 Canada, which has left a large depression. On the contrary, the gravity field over the 1 km thick permafrost (such
112 as in East Siberia) can reach 50 mGal. On the geological time scale, this field can influence the homeostasis of
113 living organisms and their speciation.

114 **6. Ocean**

115 Oceanic ice is a major control of the global climate as it screens heat and mass transfer in both vertical (across the
116 air-water interface) and horizontal directions. Warm currents, such as that of North Atlantic, move mainly along
117 the ocean surface due to thermal expansion of warm water. Glaciers on the way of warm currents can melt rapidly
118 because the ice not necessarily should reach the seafloor to stop them but can be only commensurate with the
119 surface water layer.

120 Another great challenge addressed in the PEEEX Science Plan is the ongoing change in the Arctic Ocean ice cover,
121 showing increasingly larger and longer-exposed ice-free areas (Hayes et al., 2014; Schaefer et al., 2014; Döscher
122 et al., 2014). This tendency is favorable for exploring and transportation of the Northern Siberian oil and gas
123 comprising about 25% of their total world resources (Yenikeyeff and Krysiek, 2007).

124 **7. Atmospheric ice and water**

125 Ice has very high heat capacity, specific heat of melting, and dielectric constant. The phase transition forms shields
126 of different types and scales from the gas hydrates stored in permafrost to the ice sheets that cover the entire
127 Antarctic continent and affect the global climate and evolution of life. In the global Earth system, the cryosphere

128 acts as the temperature stabilizer due to exceptional thermal inertia of ice and water occupying a great part of the
129 Earth's surface.

130 The Earth's atmosphere contains a large amount of water in the form of vapor, liquid and ice which influences the
131 planetary heat budget. Moreover, the atmosphere acts as a "thermal shield" aggregating the water molecules into
132 ice particles and thus holding back on the Earth the water, vital for the maintenance of life. The phase transitions
133 of atmospheric ice produce diverse sudden effects and various precipitation types: rain, ice, drizzle, moist snow,
134 snow, sleet, or hail. In the dry troposphere, the air cools down adiabatically at 0.98° C per 100 m. Hence,
135 condensation of originally nearly saturated water vapor starts already at a few hundred meters above the surface.
136 Along with sublimation, this leads to the heat release, which affects temperature profiles and other properties of
137 the atmosphere. Generally, condensation of pure water vapor starts at 400 - 00% supersaturation (e.g. Pruppacher
138 and Klett, 1997). However, condensation in the real atmosphere starts shortly after relative humidity exceeds 100%
139 due to the presence of solid or liquid aerosol particles acting as condensation nuclei, or even below 100% in some
140 cases (Kulmala et al., 1997). Similarly, ice nucleation is initiated upon nuclei like dust particles, biological particles
141 etc., including anthropogenic aerosol particles (e.g. Hoose and Möhler, 2010, Atkinson et al. 2013). Thus, the air
142 pollution facilitates formation of ice crystals and cloud droplets, and therefore affects the air temperature and
143 precipitation processes and its spatiotemporal variability (Rosenfeld et al. 2014). The aerosol-cloud-climate
144 interactions (see e.g. Kulmala et al., 2011), as well as air pollution-weather-climate interactions (Ding et al. 2016,
145 Petäjä et al., 2016) have been and are under rigorous investigations.

146 **8. Biota and cold**

147 A number of facts in the evolution biology would be unexplainable without cryosphere. Contrary to the common
148 thinking, the existence of life is consistent with negative temperatures. For example, photosynthesis is quite
149 possible at temperatures below 0oC (Kramer P.J., Kozłowski T.T., 1960). For billions of years, the cryosphere

150 assured the stability of environment required for development of life. The mean annual air temperatures on the
151 Earth were naturally maintained near the ice-water transition point. Ice, like water, possesses exceptional
152 thermostatic properties, and is thus a very strong stabilizer of the temperature regime, along with wide spread of
153 ice over the Earth's surface. The heat capacity of water (4.183 kJ/kg·K) is five times the mean heat capacity of soil,
154 and its volumetric heat capacity is 3300 times that of air. The high heat capacity of ice (2.06 kJ/kg·K) is comparable
155 with that of water, which makes the ice-snow-water system the Earth's largest storage of solar energy. The very
156 phase change point has additional (likewise abnormal) thermal stability: the specific heat of ice melting is five
157 times as high as in gold (332 against 66.2 kJ/kg) and is higher than, say, in mercury by a factor 28 (Melnikov and
158 Gennadinik, 2011).

159 Permafrost contains organic matter in the form of biogenic gases, products of biodegradation, or as ancient viable
160 bacteria. In Russia, the evidence of bacteria in permafrost was obtained in the latest 19th century, together with
161 finding of mammoths in Northern Siberia and during studies of soils in the Russian Far East. Novel information
162 about bacterial communities has been obtained from studies of Antarctic permafrost. Cyanobacteria, with their age
163 about 500 kyr, were discovered in the station "Vostok" ice cores at 3600 m depth (Water sampling, 2003).

164 Many microorganisms are stable against freezing and not only survive but even grow at negative temperatures.
165 More than 10% of water in organic tissues remains unfrozen at temperatures below -20°C. Many basic processes
166 in microbial, soil, geophysical, and cosmological systems are associated with cooling and ice-water phase changes.
167 Ribosome activity is often more successful at low temperatures, which may be the evidence that life could have
168 originated in cold conditions (Vlassov et al., 2005).

169 Life processes can remain active at below 0 ° C in capillaries and on ice surface in permafrost. Micro-organisms
170 can exist in media with low dimensionality (2 or less), such as liquid films and capillaries filled with super-cooled
171 water. The porous structure of the environment ensures the transport of nutrients and life wastes (Möhlmann,
172 D.T.F., 2009).

173 Even comparatively minor heating to 50-60°C leads to denaturation of proteins and can stop operation of a living
174 system, while freezing even to the near-absolute zero temperatures does not change the configuration of
175 biomolecules, so that the life functions can resume after thawing. Enzymes that regulate metabolism in organisms
176 show similar behavior (Dethier and Vilee, 2005).

177 It is still unclear how and why the metazoans originated about a billion years ago. The obviously required sufficient
178 amount of oxygen has been commonly explained by its gradual photosynthetic accumulation in the atmosphere.
179 However, origination of metazoans may have been due to the higher concentration of oxygen in cold water. The
180 origin of the skeletal organisms at the Proterozoic/Phanerozoic boundary was possibly favored by the calcium
181 carbonate release by organisms at the high oxygen availability. The oxygen enrichment in oceans at the
182 Vendian/Cambrian boundary may have resulted from glaciation, water cooling, hence, increasing gas solubility
183 and favorable water circulation patterns. Note that the first giant organisms have appeared precisely during this
184 cooling event. Furthermore, gigantism, associated with specific heat exchange, is known to be typical of polar
185 biota. Biotic and cooling events have been closely related all over the Phanerozoic. The major Jurassic and latest
186 Cretaceous extinctions occurred during the stages of cold climate (The Winters, 1979).

187 **9. Typical rates of processes**

188 Natural objects are characterized by their typical scales in time and space. The common sense suggests that
189 temporal and spatial scales change more or less similarly, which implies that typical change rates are not too
190 variable. However, this rule does not work for the cryosphere, which can be either accelerator or decelerator of
191 various processes (Melnikov et al., 2013). For instance, paleo-bacteria preserved under ice shields remain viable
192 for millions of years. Ice slow down biotic processes due to the following properties:

- 193 • low dielectric and magnetic susceptibility, which reduces electromagnetic fields;
- 194 • high stiffness, which resists mechanic impacts;

195 • low permeability, which decelerates mass transport;

196 • high heat capacity and anomalous specific heat of melting, which damp temperature variations.

197 Hence, interaction between an object and its environment in cryosphere is basically restricted to weak gravity
198 effects and very slow diffusion. The capability of ice to decelerate biological processes has practical uses: freezing
199 of organs or storage of seeds, not to mention the household refrigerators. Similar mechanisms protect the biota of
200 Antarctic paleo-lakes, such as Lake Vostok, hidden under kilometers-thick ice maintaining its sustainable thermal
201 regime (Thoma et al., 2008).

202 On the contrary, processes on the surface of ice or ice-bearing systems are often faster than usual due to formation
203 of electrolytes, chemical reactions, and accelerated transport. “Ice fingers of death” is an example of a very rapid
204 process with extremely high gradients of physical parameters in the vicinity of cryogenic objects. This phenomenon
205 (also known as brinicles, brine icicles or ice stalactites) develops in shallow water beneath the sea ice, mainly in
206 the southern high latitudes, in the form of slowly rotating jets of extremely cold and saline (and therefore heavy)
207 surface water sinking down to the ocean bottom (Martin, 1974). The point is that the formation of sea ice at very
208 low air temperatures is associated with the decomposition of calcium bicarbonate into CO_2 and CaCO_3 . Then CO_2
209 releases into the air, whereas slowly rotating jets of the cold dense surface water sinks down to the seafloor and
210 propagates downslope leaving frozen footprints with entrapped bottom-dwelling sea animals.

211 Cryogenic phenomena developing at various rates are presented in the log-log space/time coordinates in Fig. 4
212 showing extremely wide range of scales. Some objects, such as paleo-bacteria survived in ice over half a million
213 years or “ice fingers of death”, fall beyond the scope of any particular science. To emphasize their multidisciplinary
214 nature, these objects are marked by two colors indicating the relevant pairs of disciplines.

215 **10. Cryodiversity and its role in the Earth system**

216 The Earth's cryosphere spreads from the oceanic depths reaching ~5 km to the boundary between the atmosphere
217 and outer space (called Karman line) at ~100 kilometers above the sea level. The cryospheres of the Earth and other
218 planets with their satellites actually comprise the cryosphere of the Sun System with its radius ~100 au, where the
219 density of solar wind (radiated plasma) is high. Thus, the terrestrial cryosphere is an object among a diversity of
220 planets and other large objects in the Solar System, which differ in their distances to the Sun; rotation rates and
221 orbiting paths with the respective diurnal and annual cycles; chemical compositions; and presence or absence of
222 the atmospheres. Some planets of the Solar System and their satellites store great amounts of ice, much more than
223 the Earth, both in absolute values and relative to the mass of the respective celestial bodies. Recent extraterrestrial
224 missions have demonstrated impressive cryogenic processes and phenomena on various planets and satellites
225 (Komarov and Isaev, 2010).

226 **11. Cryogenic objects in the Solar system**

227 Historically, the Earth's cryosphere was investigated in the scope of exploration of the high-latitude or mountain
228 regions characterized by cold climate, large stocks of ice and snow, and strong thermal effects of the H₂O phase
229 transitions. Those investigations were basically performed in the conceptual framework of geographical disciplines,
230 such as glaciology or permafrost science. The first priority belonged to description, quantification and inventory of
231 phenomena, with comparatively little attention to concrete features of the background physical, chemical and
232 biological processes. Such approach, although motivated by immediate practical purposes, limits our knowledge
233 of the cryosphere basically to its present state and geological past.

234 In view of the ongoing climate change and massive anthropogenic deterioration of the environment, it becomes
235 critical to be able to predict the near future of the Earth's cryosphere in its capacity as the major stabilizer of the
236 Earth climate and environment. To this end, a leap is needed from the currently dominating descriptive approach

237 to the process-based approach unifying investigation of the totality of cryogenic phenomena from molecular to
238 planetary scales (see Figure 4) in an interdisciplinary science, *cryology*, armed with all necessary methods of
239 geosciences, physics, chemistry and biology.

240 **12. Concluding remarks**

241 Ice has very high heat capacity, specific heat of melting, and dielectric constant. The phase transition forms shields
242 of different types and scales from the gas hydrates stored in permafrost to the ice sheets that cover the entire
243 Antarctic continent and affect the global climate and evolution of life. In the global Earth system, the cryosphere
244 acts as the temperature stabilizer due to exceptional thermal inertia of ice and water occupying a great part of the
245 Earth's surface.

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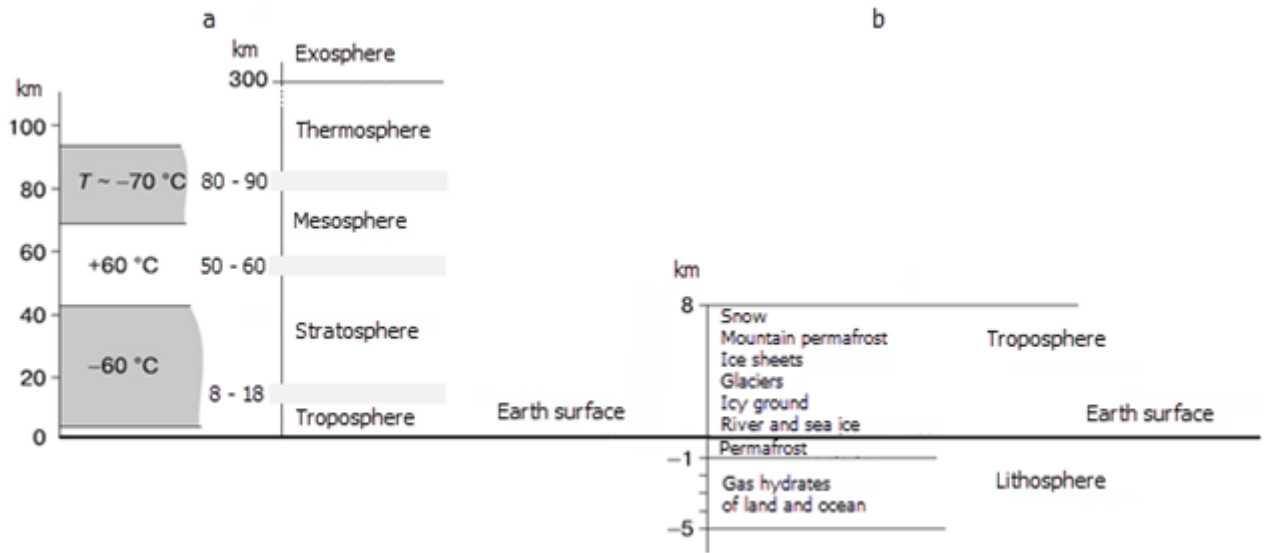
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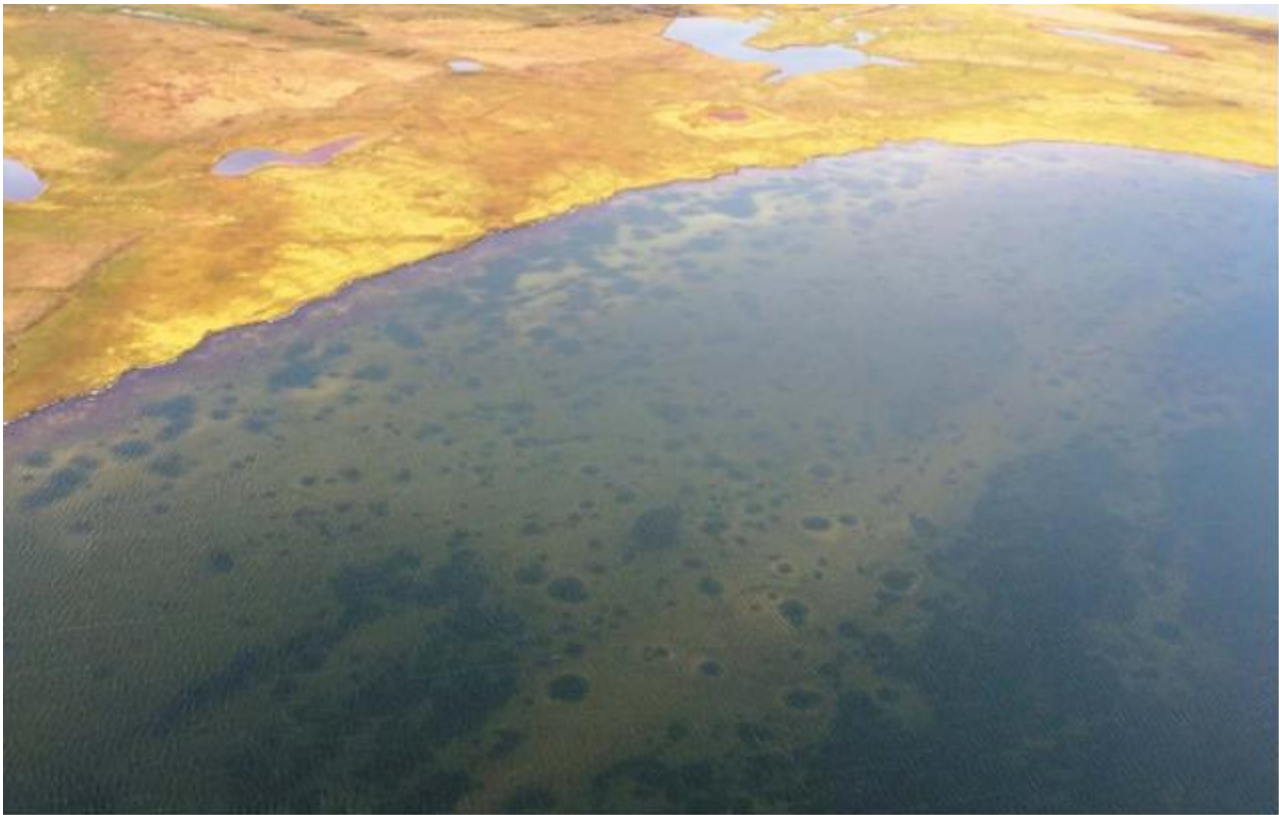
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427 Figure 1: The Earth's cryosphere: from (a) the uppermost atmosphere, to (b) 5 km below Earth's surface, where

428 ice exists as hydrates at positive temperatures.



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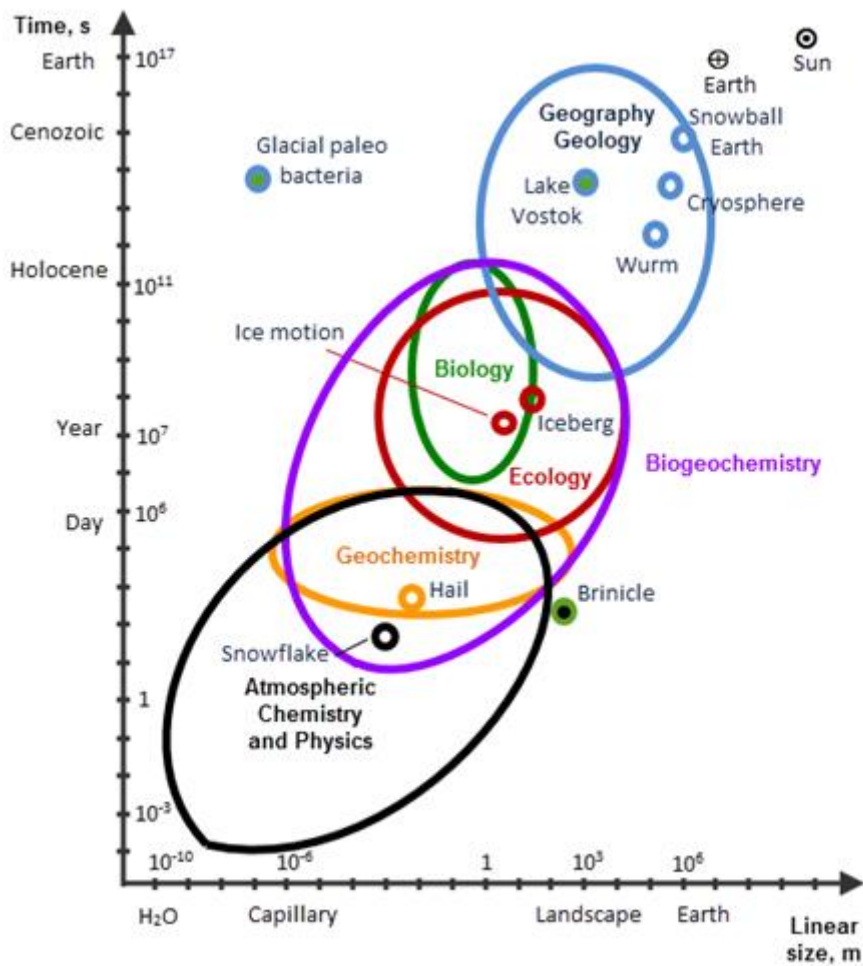
430 Figure 2: Photo of a lake on the Yamal Peninsula showing numerous craters from gas explosions on the lake bottom

431 (Bogoyavlensky, 2015).



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433 Figure 3: New crater produced by gas explosion on the Yamal Peninsula in 2014 (Leybman et al., 2015; photograph
434 by Vladimir Pushkarev borrowed from “Siberian Times”).



435

436 Figure 4: Log-log space-time diagram of cryogenic phenomena of different nature. Large ellipses or circles with

437 different colors identify various sciences. Small circles show various objects of the cryosphere. The Earth and the

438 Sun are shown to illustrate the range of scales.