



1 **Predicting the glass transition temperature and viscosity of secondary**
2 **organic material using molecular composition**

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14



15 **Abstract:**

16 Secondary organic aerosols (SOA) account for a large fraction of submicron particles in the
17 atmosphere. SOA can occur in amorphous solid or semi-solid phase states depending on
18 chemical composition, relative humidity (RH), and temperature. The phase transition between
19 amorphous solid and semi-solid states occurs at the glass transition temperature (T_g). We have
20 recently developed a method to estimate T_g of pure compounds containing carbon, hydrogen, and
21 oxygen atoms (CHO compounds) with molar mass less than 450 g mol^{-1} based on their molar
22 mass and atomic O:C ratio. In this study, we refine and extend this method for CH and CHO
23 compounds with molar mass up to $\sim 1100 \text{ g mol}^{-1}$ using the number of carbon, hydrogen, and
24 oxygen atoms. We predict viscosity from the T_g -scaled Arrhenius plot of fragility (viscosity vs.
25 T_g/T) as a function of the fragility parameter D . We compiled D values of organic compounds
26 from literature, and found that D approaches a lower limit of ~ 10 (+/- 1.7) as the molar mass
27 increases. We estimated viscosity of α -pinene and isoprene SOA as a function of RH by
28 accounting for hygroscopic growth of SOA and applying the Gordon-Taylor mixing rule,
29 reproducing previously published experimental measurements very well. Sensitivity studies were
30 conducted to evaluate impacts of T_g , D , hygroscopicity parameter (κ), and the Gordon-Taylor
31 constant on viscosity predictions. Viscosity of toluene SOA was predicted using the elemental
32 composition obtained by high-resolution mass spectrometry (HRMS), resulting in a good
33 agreement with the measured viscosity. We also estimated viscosity of biomass burning particles
34 using the chemical composition measured by HRMS with two different ionization techniques:
35 electrospray ionization (ESI) and atmospheric pressure photoionization (APPI). Due to
36 differences in detected organic compounds and signal intensity, predicted viscosities at low RH
37 based on ESI and APPI measurements differ by 2-5 orders of magnitude. Complementary



38 measurements of viscosity and chemical composition are desired to further constrain RH-
39 dependent viscosity in future studies.

40

41 **1. Introduction**

42 Secondary organic aerosols (SOA) account for a large fraction of submicron particles in the
43 atmosphere and they play an important role in climate, air quality and public health (Goldstein
44 and Galbally, 2007; Jimenez et al., 2009). Traditionally, SOA particles were assumed to be
45 liquid with dynamic viscosity η below 10^2 Pa s, but a number of recent studies have shown that
46 they can also adopt amorphous semi-solid ($10^2 \leq \eta \leq 10^{12}$ Pa s), or glassy solid ($\eta > 10^{12}$ Pa s)
47 states, depending on chemical composition, relative humidity (RH) and temperature (Zobrist et
48 al., 2008; Mikhailov et al., 2009). Ambient and laboratory-generated SOA particles have been
49 observed to bounce off the smooth hard surface of an inertial impactor at low RH, implying a
50 non-liquid state (Virtanen et al., 2010; Saukko et al., 2012; Bateman et al., 2015; Jain and
51 Petrucci, 2015), whereas predominantly biogenic SOA particles in the Amazon basin did not
52 bounce off the impactor surface at high RH, implying they are primarily liquid (Bateman et al.,
53 2016). Upon dilution or heating, SOA particles were observed to evaporate unexpectedly slowly,
54 consistent with their (semi-)solid phases (Cappa and Wilson, 2011; Vaden et al., 2011; Yli-Juuti
55 et al., 2017). Direct measurements of viscosity of SOA bulk material derived from oxidation of
56 α -pinene (Renbaum-Wolff et al., 2013; Zhang et al., 2015; Hosny et al., 2016), limonene (Hinks
57 et al., 2016), isoprene (Song et al., 2015), and toluene (Song et al., 2016a) have confirmed that
58 SOA particles adopt a wide range of viscosities.

59 The particle phase state has been shown to affect gas uptake and chemical transformation of
60 organic compounds due to kinetic limitations of bulk diffusion (Shiraiwa et al., 2011; Abbatt et



61 al., 2012; Kuwata and Martin, 2012; Zhou et al., 2013; Slade and Knopf, 2014; Arangio et al.,
62 2015; Davies and Wilson, 2015; Wang et al., 2015; Berkemeier et al., 2016; Marshall et al.,
63 2016). Molecular motion can be hindered in a highly viscous matrix, slowing down
64 photochemical reactions in particles (Lignell et al., 2014; Hinks et al., 2016). Partitioning of
65 semi-volatile compounds into viscous particles may result in kinetically-limited growth in
66 contrast to quasi-equilibrium growth (Perraud et al., 2012; Shiraiwa and Seinfeld, 2012; Shiraiwa
67 et al., 2013; Booth et al., 2014; Zaveri et al., 2014; Liu et al., 2016). Water diffusion can be still
68 fast even in an amorphous solid matrix under room temperature, but it can be hindered
69 significantly under low temperatures (Mikhailov et al., 2009; Zobrist et al., 2011; Bones et al.,
70 2012; Berkemeier et al., 2014; Price et al., 2014), affecting homogeneous vs. heterogeneous ice
71 nucleation pathways (Murray et al., 2010; Wang et al., 2012; Baustian et al., 2013; Schill and
72 Tolbert, 2013; Berkemeier et al., 2014; Schill et al., 2014; Lienhard et al., 2015; Ignatius et al.,
73 2016). Despite the substantial implications of the SOA particle phase state, its effects on gas-
74 particle interactions have not yet been considered explicitly in current climate and air quality
75 models (Shrivastava et al., 2017).

76 Group contribution methods have been used to predict the viscosities of pure compounds.
77 Song et al. (2016) showed that estimations from group contribution approaches combined with
78 either nonideal or ideal mixing reproduced the RH-dependent trends particularly well for the
79 alcohol, di-, and tricarboxylic acid systems with viscosity of up to 10^4 Pa s. By contrast, model
80 calculations overestimated the viscosity of more viscous compounds including mono-, di-, and
81 trisaccharides by many orders of magnitude (Song et al., 2016b). A recent study compiled
82 viscosity of organic compounds with atmospherically relevant functional groups, investigating
83 the influence of the number and location of functional groups on viscosity (Rothfuss and Petters,



84 2017). Information on molecular structures, however, is often difficult to obtain in laboratory or
85 field measurements on atmospheric aerosols.

86 Particle phase state can be characterized by a glass transition temperature (T_g), which is a
87 characteristic temperature representing a non-equilibrium phase transition from a glassy solid
88 state to a semi-solid state as the temperature increases (Koop et al., 2011). Recently, we have
89 developed a method to estimate T_g of pure organic compounds comprised of carbon, hydrogen,
90 and oxygen (CHO compounds) with molar mass less than 450 g mol^{-1} based on their molar mass
91 and atomic O:C ratio (Shiraiwa et al., 2017). It has been applied successfully in a global
92 chemistry climate model to predict T_g and the phase state of atmospheric SOA. Global
93 simulations indicate that SOA particles are mostly liquid or semi-solid in the planetary boundary
94 layer, while they should be glassy in the middle and upper troposphere (Shiraiwa et al., 2017). A
95 very recent study provided a consistent result, suggesting that mixing timescales of organic
96 molecules within SOA are often $< 1 \text{ h}$ in a global planetary boundary layer (Maclean et al.,
97 2017).

98 It has been shown that SOA particles contain oligomeric compounds with molar masses
99 higher than 450 g mol^{-1} (Gao et al., 2004; Tolocka et al., 2004; Nizkorodov et al., 2011; Nozière
100 et al., 2015), which makes the previously developed parameterization incomplete. In this study,
101 we extend the parameterization of T_g to higher molar mass compounds, and apply it to high-
102 resolution mass spectrometry data for toluene SOA and biomass burning particles. The
103 Arrhenius approach and the Gordon-Taylor mixing rules were applied to estimate viscosity of
104 SOA bulk materials to compare with the literature reported viscosity measurements. This method
105 will be useful for estimations of viscosity of organic particles, for which high-resolution mass



106 spectra are available. It can also be applied in global or regional models to evaluate impacts of
107 the particle phase state on the role of SOA in climate and air quality.

108

109 **2. Parameterization development**

110 **2.1 Glass transition temperature**

111 Figure 1a shows the dependence of T_g on the molar mass (M) of organic compounds. Solid
112 markers represent measured T_g of 258 CHO compounds (Koop et al., 2011; Dette et al., 2014;
113 Rothfuss and Petters, 2017), while open markers represent 654 CHO compounds in SOA
114 (Shiraiwa et al., 2014). Markers are color-coded by atomic O:C ratio. Their melting points (T_m)
115 were estimated by the Estimation Programs Interface (EPI) Suite software version 4.1 (US-EPA,
116 2012) and their T_g were estimated using the Boyer-Kauzmann rule: $T_g = g \cdot T_m$ with $g = 0.7$ (Koop
117 et al., 2011; Shiraiwa et al., 2017). A subset of data shown in Figure 1 was originally published
118 in Shiraiwa et al. (2017) for compounds with $M < 450 \text{ g mol}^{-1}$. This version of the figure has
119 been updated to include a number of experimentally measured T_g values of larger compounds
120 with M up to 1153 g mol^{-1} , including aliphatic compounds containing OH and/or COOH groups.
121 Specifically, data for 76 aliphatic alcohols, 39 carbohydrates and their derivatives, 4 carboxylic
122 acids, and 4 hydroxy acids, as compiled by Rothfuss and Petters (2017), have been added to
123 Figure 1. These updates are critical for reliable parametrization of T_g based on M . When M
124 increases above $\sim 500 \text{ g mol}^{-1}$, the slope of T_g decreases, making it challenging to extrapolate the
125 low- M data from the original Shiraiwa et al. (2017) study to higher M values. When M increases
126 to $\sim 1000 \text{ g mol}^{-1}$, the corresponding T_g appears to level at around 420 K.

127 Such dependence on M has been described for polymers with the Fox-Flory equation:

128 $T_g(M) = T_{g,\infty} - \frac{K_m}{M}$ (Fox Jr and Flory, 1950), where K_m is a constant and $T_{g,\infty}$ is the asymptotic



129 value of T_g specific to the polymer. We conducted a literature search and found that most of the
130 reported $T_{g,\infty}$ values fell below ~ 500 K (Fox Jr and Flory, 1950; Onder et al., 1972; Montserrat
131 and Colomer, 1984; Polymer handbook, 1999; Papadopoulos et al., 2004; Matsushima et al.,
132 2017). The Fox-Flory equation works very well for high molar mass compounds and is also
133 generally applicable to smaller compounds (Koop et al., 2011), as supported by an approximately
134 linear dependence of T_g on the inverse molar mass in Fig. A1. Figure 1b plots the values of T_g as
135 a function of the atomic O:C ratio of organic molecules. Figures 1a and 1b clearly demonstrate
136 that T_g depends primarily on the molar mass with a weak dependence on the atomic O:C ratio.

137 A parameterization for T_g calculation based on the molar mass and atomic O:C ratio was
138 developed in our recent work, which is applicable to CH and CHO compounds with $M < 450$ g
139 mol^{-1} (Shiraiwa et al., 2017):

$$140 \quad T_g = A + BM + C^2M^2 + D(\text{O:C}) + E M(\text{O:C}) \quad (1)$$

141 where $A = -21.57 (\pm 13.47)$ [K], $B = 1.51 (\pm 0.14)$ [K mol g^{-1}], $C = -1.7 \times 10^{-3} (\pm 3.0 \times 10^{-4})$ [K
142 $\text{mol}^2 \text{g}^{-2}$], $D = 131.4 (\pm 16.01)$ [K] and $E = -0.25 (\pm 0.085)$ [K mol g^{-1}], respectively. These values
143 were obtained by fitting the measured T_g of 179 CH and CHO compounds with $M < 450$ g mol^{-1}
144 with multi-linear least squares analysis. Note that application of Eq. (1) may provide
145 unreasonable T_g values for compounds with $M > 500$ g mol^{-1} because it does not account for the
146 strong curvature in the T_g vs. M dependence shown in Figure 1a.

147 In this study we have developed an improved parameterization to predict T_g of CH and
148 CHO compounds using the number of carbon (n_C), hydrogen (n_H), and oxygen (n_O) that can also
149 be applied to higher molar mass compounds. Motivated by a good correlation between T_g and
150 volatility (Fig. 1a in Shiraiwa et al., (2017)), we use an equation with a similar formulation to the



151 equation used to predict the saturation mass concentration or volatility (Donahue et al., 2011; Li
152 et al., 2016):

$$153 \quad T_g = (n_C^0 + \ln(n_C)) b_C + \ln(n_H) b_H + \ln(n_C) \ln(n_H) b_{CH} + \ln(n_O) b_O + \ln(n_C) \ln(n_O) b_{CO} \quad (2)$$

154 where n_C^0 is the reference carbon number, b_C , b_H and b_O denote the contribution of each atom to
155 T_g , and b_{CH} and b_{CO} are coefficients that reflect contributions from carbon-hydrogen and carbon-
156 oxygen bonds, respectively. These values were obtained by fitting the measured T_g of 42 CH
157 compounds and 258 CHO compounds with multi-linear least squares analysis with 68%
158 prediction and confidence intervals. The best-fit parameters are summarized in Table 1.

159 Note that the evaluation dataset used to derive Eq. (2) contains CH compounds with $M <$
160 260 g mol^{-1} (see Fig. A2a for comparison of measured and predicted T_g). Thus, the application of
161 Eq. (2) to higher molar mass CH compounds may require further investigations when measured
162 T_g for higher molar mass CH compounds becomes available. Figure 1c shows that the T_g values
163 predicted using Eq. (2) are in good agreement with the T_g values measured in experiments (see
164 also Fig. A2b) or estimated by the Boyer-Kauzmann rule as indicated by the high correlation
165 coefficient of 0.95. T_g of individual compounds can be predicted within $\pm 21 \text{ K}$ as indicated by
166 the prediction band (dotted lines in Fig. 1c); however, this uncertainty may be much smaller for
167 multicomponent SOA mixtures under ideal mixing conditions as indicated in the confidence
168 band (dashed lines, almost overlapping with the 1:1 line).

169 These results are noteworthy given that the parameterization (Eq. 2) does not consider
170 either explicit molecular structures or functional groups. Previous studies have shown that T_g can
171 be especially sensitive to the number of OH groups, which interact strongly through hydrogen
172 bonding. For example, Nakanishi et al., (2011) found a direct relationship between T_g and the
173 number of hydroxyl groups in a molecule for sugar alcohols; T_g increases as the number of OH



174 groups increases. They reported that the correlation between T_g and the number of OH groups
175 was much stronger than the correlation between T_g and the number of carbons in a molecule.
176 Such a trend is implicitly included in Eq. (1) and (2), which contain the O:C ratio and number of
177 oxygen atoms as parameters, respectively. Recently, Rothfuss and Petters (2017) showed an
178 approximately linear relationship between the number of OH groups and T_g for compounds with
179 up to eight OH groups. Grayson et al. (2017) showed that addition of hydroxyl functional groups
180 increases viscosity, a conclusion supported by both the experimental data and quantitative
181 structure-property relationship model. The correlation between T_g and the number of carbon
182 atoms is consistent with the free volume theory, while the correlation with the number of OH
183 groups is more consistent with the topological constraint theory (Nakanishi and Nozaki, 2011;
184 van der Sman, 2013). Future experiments targeting more comprehensive T_g data, especially for
185 higher molar mass compounds, would lead to further refinements of our T_g parameterizations.

186 Comparing Eq. (1) and (2), the two parameterizations give similar performance for
187 compounds with $M < 450 \text{ g mol}^{-1}$ as shown in Fig. A2c. The statistical measures of correlation
188 coefficient (R), mean bias (MB), and root mean square error (RMSE) are 0.93, -6.45 K , and
189 25.64 K , respectively, for the performance of Eq. (1), while for Eq. (2), they are 0.95, 3.15 K ,
190 and 21.11 K , respectively. It should be noted again that Eq. (1) cannot be used to predict T_g for
191 compounds with $M > 450 \text{ g mol}^{-1}$. For example, T_g of stachyose ($M = 667 \text{ g mol}^{-1}$) predicted by
192 Eq. (1) is 198 K , while that by Eq. (2) is 394 K , which agrees much better with the measured
193 mean T_g of 396 K (Rothfuss and Petters, 2017). Eq. (2) is more flexible than Eq. (1) and can be
194 potentially expanded to include compounds containing hetero-atoms (e.g., nitrogen or sulfur),
195 once substantial sets of experimental values of T_g for such compounds become available.
196 Regarding the application in air quality and climate models, Eq. (1) can be applied in the



197 volatility basis set (VBS) (Donahue et al., 2006; Donahue et al., 2011) and the molecular
 198 corridor approach (Shiraiwa et al., 2014; Li et al., 2016) to predict the T_g of SOA particles
 199 (Shiraiwa et al., 2017), while the new parameterization may be suitable for coupling with the
 200 statistical oxidation model which characterizes the SOA evolution as a function of n_C and n_O
 201 (Cappa and Wilson, 2012; Jathar et al., 2015).

202 These parameterizations (Eqs. 1, 2) calculate T_g based on the elemental composition of
 203 organic compounds. SOA particles contain a number of organic compounds as well as a variable
 204 amount of liquid water. Estimations of T_g for SOA-water mixtures were discussed by Shiraiwa et
 205 al. (2017), who applied the Gordon-Taylor equation validated for a wide range of mixtures of
 206 organics, polymer, and water (Roos, 1993; Hancock and Zografí, 1994; Zobrist et al., 2008;
 207 Dette et al., 2014; Dette and Koop, 2015). Briefly, T_g of mixtures of SOA compounds under dry
 208 conditions ($T_{g,org}$) were calculated assuming the Gordon-Taylor constant (k_{GT}) of 1 (Dette et al.,
 209 2014): $T_{g,org} = \sum_i w_i T_{g,i}$, where w_i is the mass fraction of organic compound i , which can be
 210 derived using mass concentrations of SOA products. Under humid conditions, SOA particles
 211 take up water by hygroscopic growth in response to RH. The Gordon-Taylor equation can also
 212 be applied to calculate T_g of organic-water mixtures considering the mass fraction of organics
 213 (w_{org}) in SOA particles (Koop et al., 2011):

$$214 \quad T_g(w_{org}) = \frac{(1-w_{org})T_{g,w} + \frac{1}{k_{GT}}w_{org}T_{g,org}}{(1-w_{org}) + \frac{1}{k_{GT}}w_{org}} \quad (3)$$

215 w_{org} can be calculated using the mass concentrations of water (m_{H_2O}) and SOA (m_{SOA}) as $w_{org} =$
 216 $m_{SOA} / (m_{SOA} + m_{H_2O})$. m_{H_2O} can be estimated using the effective hygroscopicity parameter (κ)
 217 (Petters and Kreidenweis, 2007):

$$218 \quad m_{H_2O} = \frac{\kappa \rho_w m_{SOA}}{\rho_{SOA} \left(\frac{1}{a_w} - 1 \right)} \quad (4)$$



219 The density of water (ρ_w) is 1 g cm^{-3} , the density of SOA particles (ρ_{SOA}) is assumed to be 1.2 g
220 cm^{-3} (Kuwata et al., 2012), m_{SOA} is the total mass concentrations of SOA, and a_w is the water
221 activity calculated as $a_w = \text{RH}/100$.

222

223 2.2 Viscosity

224 Temperature dependence of viscosity (η) was predicted using the modified Vogel-
225 Tamman-Fulcher (VTF) equation (Angell, 1991):

$$226 \quad \eta = \eta_{\infty} e^{\frac{T_0 D}{T - T_0}} \quad (3)$$

227 where η_{∞} is viscosity at infinite temperature; T_0 is the Vogel temperature; T is the ambient
228 temperature. The fragility parameter, D , characterizes how rapidly the dynamics of a material
229 slow down as T approaches T_g , reflecting to what degree the temperature dependence of the
230 viscosity deviates from Arrhenius behavior. When T is close to T_g ($T_g/T \approx 1$), smaller D values
231 indicate that viscosity is sensitive to temperature change (fragile behavior); while larger D values
232 indicate that viscosity is less sensitive to temperature change (strong or Arrhenius behavior).

233 Assuming $\eta_{\infty} = 10^{-5} \text{ Pa s}$ (Angell, 1991):

$$234 \quad \log \eta = -5 + 0.434 \frac{T_0 D}{T - T_0} \quad (4)$$

235 When $T = T_g$, $\eta = 10^{12} \text{ Pa s}$, which leads to (Angell, 1991; Angell, 2002):

$$236 \quad T_0 = \frac{39.17 T_g}{D + 39.17} \quad (5)$$

237 As can be seen in Eq. (5), both T_g and D are required to calculate η from Eq. (4) at a given
238 temperature.

239 Figure 2 shows the T_g -scaled Arrhenius plot of fragility (viscosity versus T_g/T) referred to
240 as an Angell plot (Angell, 1995). D values of organic compounds are typically in the range of
241 ~ 5 – 30 (Angell, 1997). To estimate D values that could be applied to SOA compounds, we



242 compiled measured fragility values. Fragility was often measured in the form of the fragility
243 steepness index (m), which represents the slope of the Arrhenius plot at the point where $T = T_g$
244 (Boehmer et al., 1993). Compounds with lower m exhibit higher D values, indicating stronger
245 glass formers. The measured m of 95 organic compounds are included in the Supplement. m can
246 be converted to D using the following equation (see the full derivation of this equation in
247 Appendix A):

$$248 \quad D = \frac{665.89}{m-17} \quad (6)$$

249 Figure 3 shows the measured D as a function of (a) molar mass and (b) atomic O:C ratio
250 of organic molecules. The molar mass exerts a stronger effect on fragility, while there is little
251 dependence of D on the O:C ratio. As molar mass increases, D approaches a lower limit of 10.3
252 (± 1.7), consistent with the value of 10 used in our recent study (Shiraiwa et al., 2017). To
253 evaluate the impact of the variations of D on viscosity prediction, sensitivity calculations were
254 conducted as described in Sect. 3.

255 Besides the VTF equation, another commonly used equation for describing the
256 temperature dependence of viscosity is the Williams-Landel-Ferry (WLF) equation: $\log \frac{\eta(T)}{\eta(T_g)} =$
257 $\frac{-C_1(T-T_g)}{C_2+(T-T_g)}$, where empirical parameters C_1 and C_2 are adopted as 17.44 and 51.6 K, respectively
258 (Williams et al., 1955; Schill and Tolbert, 2013; Wang et al., 2015). The two equations are
259 mathematically equivalent and the WLF and VTF parameters are related through $C_1 =$
260 $\frac{DT_0}{2.303(T_g-T_0)}$ and $C_2 = T_g - T_0$. It is shown that, compared to the WLF equation, the VTF
261 parameterization does not rely heavily on the free volume assumption and the Vogel temperature
262 T_0 has a more profound physical meaning that relates both thermodynamic and kinetic concepts
263 (O'Connell and McKenna, 1999; Huang and McKenna, 2001). The calculations of viscosity in



264 this study are based mainly on the VTF equation and the difference between calculated results
265 from the two equations will be briefly discussed in the following section.

266

267 **3. Comparison of predicted viscosity with measurements**

268 **3.1. SOA formed from α -pinene and isoprene**

269 The purpose of this section is to demonstrate that viscosity of SOA material can be
270 predicted over a broad range of RH values from four parameters: T_g of dry SOA ($T_{g,org}$), fragility
271 (D), hygroscopicity (κ), and the Gordon-Taylor constant for mixing SOA and water (k_{GT}).
272 Viscosity of α -pinene SOA has been measured as a function of RH by several groups using
273 multiple experimental techniques as shown in Fig. 4(a) (Abramson et al., 2013; Renbaum-Wolff
274 et al., 2013; Kidd et al., 2014; Pajunoja et al., 2014; Bateman et al., 2015; Zhang et al., 2015;
275 Grayson et al., 2016). These datasets suggest that viscosity of α -pinene SOA is very high ($\sim 10^8$
276 Pa s) at low RH and decreases with an increase in RH reaching a value of ~ 10 Pa s at 80% RH.
277 As can be seen in Fig. 4(b), isoprene SOA is less viscous with $\eta < 10^6$ Pa s even under dry
278 conditions, undergoing a phase transition from a semi-solid phase to a liquid phase at $\sim 55\%$ RH
279 (Bateman et al., 2015; Song et al., 2015).

280 The solid lines with the shaded areas in Figure 4 are viscosity values predicted using
281 $T_{g,org}$, D , κ , k_{GT} . $T_{g,org}$ values were adopted by Berkemeier et al. (2014) who estimated $T_{g,org}$ with
282 the Boyer-Kauzmann rule using the melting point of representative SOA oxidation products.
283 Note that Eq. (1) or (2) were not used to estimate $T_{g,org}$, which should be done in future studies
284 by obtaining their elemental composition using high resolution mass spectrometry. For α -pinene,
285 $T_{g,org}$ was assumed to be 278 K corresponding to an O:C ratio of 0.5 (Berkemeier et al., 2014),
286 which is a typical O:C ratio of α -pinene SOA (Aiken et al., 2008; Chen et al., 2011; Putman et



287 al., 2012). The wide range of viscosities reported for α -pinene SOA may indicate that the O:C
288 values may be different in different experiments. $T_{g,org}$ selected for isoprene SOA was 255 K,
289 corresponding to the O:C ratio of 0.6. Although no measurements of the O:C ratio for the
290 experimental isoprene SOA data were reported, Song et al. (2015) estimated O:C of 0.64-1.1
291 based on literature values. As O:C ratios are useful in estimating $T_{g,org}$, we encourage the
292 measurement of the O:C ratio of SOA when conducting viscosity measurements.

293 For both α -pinene and isoprene SOA, D was set to 10 based on the analysis presented in
294 Fig. 3(a). κ was set to 0.1 based on field and laboratory measurements (Gunthe et al., 2009;
295 Lambe et al., 2011b; Pajunoja et al., 2014; Petters et al., 2017) and k_{GT} was assumed to be 2.5
296 (Zobrist et al., 2008; Koop et al., 2011). Using these parameters, the predicted viscosities match
297 well the magnitude and the RH-dependence of the measured viscosity of α -pinene and isoprene
298 SOA. Figure 4 also shows predicted viscosities (dotted lines) using the WLF equation, which
299 shows similar values as the VTF equation, but slightly underestimates the viscosity of α -pinene
300 SOA at low RH and overestimates the viscosity of isoprene SOA at high RH.

301 Sensitivity studies were conducted to examine the effects of $T_{g,org}$, D , κ and k_{GT} , on the
302 calculated viscosity. In these studies, $T_{g,org}$ of α -pinene and isoprene SOA were varied within 229
303 - 328 K and 255 - 316 K, respectively, representing $T_{g,org}$ of different oxidation states
304 (Berkemeier et al., 2014). D was varied between 5 and 30, which is the range characteristic for
305 organic compounds (see Fig. 3a). κ of 0.05 - 0.15 were used for α -pinene and isoprene SOA
306 (Lambe et al., 2011b; Pajunoja et al., 2015). For the Gordon-Taylor constant, values of 2.5 ± 1.5
307 were considered (Zobrist et al., 2008; Koop et al., 2011; Dette et al., 2014; Dette and Koop,
308 2015).



309 The effect of varying each parameter on the calculated viscosity of α -pinene SOA is
310 illustrated in Fig. 5. Variations of ± 50 K in $T_{g,org}$ result in 3-6 orders of magnitude difference in
311 calculated values at dry conditions, indicating that $T_{g,org}$ is a critical parameter for viscosity
312 estimations. Decreasing D from 10 to 5 led to a decrease of calculated values by more than one
313 order of magnitude. The calculated results were within the upper limit of measurements when
314 increasing D from 10 to 20, and the predicted values were only slightly enhanced when further
315 increasing D from 20 to 30. Calculated values with variations in κ from 0.05 to 0.15 and k_{GT}
316 from 1.0 to 4.0 were all within the measured ranges.

317 For isoprene SOA, an increase of $T_{g,org}$ to 287 K, which represents a higher oxidation
318 state (Berkemeier et al., 2014), led to calculated values to be several orders of magnitude higher
319 than the upper limit of measurements (Fig. 6a). When $T_{g,org}$ reaches 316 K, isoprene SOA can
320 occur as a solid for RH lower than $\sim 40\%$. Compared to α -pinene SOA, a variation in D has a
321 larger effect on the calculated viscosity (Fig. 6b). For a range of 5 - 30 for D , calculations with
322 the D value of 10 agreed well with the measurements, while other D values resulted in calculated
323 viscosity outside of the measured ranges. Figures 6c and 6d show that decreasing κ and k_{GT}
324 below the reference values, the predictions overestimate the measured η by one or two orders of
325 magnitude. The latter is most evident at RH $> 60\%$, where the calculated values were higher than
326 the upper limit of measurements. Modeling results with κ and k_{GT} increasing to 0.15 and 4.0,
327 respectively, were within the lower limit of measurements.

328 The above comparison between the measured and predicted viscosity demonstrates that
329 the method described in this study can reproduce reasonably well the measured RH-dependent
330 viscosity of SOA formed from α -pinene and isoprene. The sensitivity calculations showed that
331 $T_{g,org}$ contributed the most to the uncertainty in the viscosity estimates. Previous studies have



332 shown that the experimental conditions such as particle mass concentrations (Grayson et al.,
333 2016) and RH upon SOA formation (Kidd et al., 2014) can impact chemical composition of SOA
334 and hence the viscosity. Further efforts to constrain the uncertainties are needed both in
335 experiments and parameterizations.

336

337 **3.2. SOA formed from toluene**

338 In this and the following sections, we examine the feasibility of calculating the value of
339 $T_{g,org}$ from mass spectrometry data on SOA. Hinks et al. (2017) measured the elemental
340 composition of toluene SOA using nanospray desorption electrospray ionization high-resolution
341 mass spectrometry (nano-DESI-HRMS) (Roach et al., 2010a, b). Toluene SOA were formed by
342 OH photooxidation in an aerosol smog chamber at <2% RH (mass loading = $23 \mu\text{g m}^{-3}$) and 75%
343 RH (mass loading = $8 \mu\text{g m}^{-3}$) to investigate the effect of RH on the chemical composition of
344 toluene SOA formed under low- NO_x conditions. Measurements revealed a significant reduction
345 in the fraction of oligomers present in toluene SOA generated under high RH conditions
346 compared to SOA generated under low RH conditions (Hinks et al., 2017). The detected molar
347 mass of individual oxidation products spanned a range of 102 - 570 g mol^{-1} at high RH, which
348 increased up to 726 g mol^{-1} at low RH.

349 Figure 7(a) shows the interdependence of glass transition temperature, volatility, and
350 molar mass. Glass transition temperatures of detected compounds were calculated using Eq. (2).
351 Saturation mass concentrations or volatilities of detected compounds were estimated from the
352 elemental composition by using the parameterization of Li et al. (2016). The analysis is based on
353 the molecular corridor approach—a two-dimensional framework of volatility and molar mass of
354 SOA components constrained by boundary lines of low and high atomic O:C ratio,



355 corresponding to *n*-alkanes (C_nH_{2n+2} , O:C = 0) and sugar alcohols ($C_nH_{2n+2}O_n$, O:C = 1),
356 respectively (Shiraiwa et al., 2014; Li et al., 2016). The toluene SOA constituents are well
357 constrained by the molecular corridor and T_g are higher for compounds with higher molar mass
358 and lower volatility.

359 Eq. (1) was used to calculate T_g for individual compounds with $M < 450 \text{ g mol}^{-1}$, while
360 excluding compounds with molar mass higher than 450 g mol^{-1} . This approach was deemed
361 reasonable as such high molar mass compounds account for < 10% of all toluene SOA products
362 formed at low RH, and for < 2% formed at high RH. Eq. (2) was used to calculate T_g for all the
363 detected compounds. T_g of dry toluene SOA ($T_{g,\text{org}}$) was then computed using the Gordon-Taylor
364 approach with $k_{GT} = 1$ (Sect. 2.1). The relative mass concentrations of individual components
365 were assumed to be proportional to their relative abundance in the nano-DESI-HRMS spectrum.
366 This assumption has a number of caveats (Bateman et al., 2012; Nguyen et al., 2013), and as we
367 will see below, it results in deviations between the predicted and measured viscosity. Table 2
368 summarizes the results of such calculations, showing that the $T_{g,\text{org}}$ by Eq. (1) – excluding high
369 molar mass compounds – is about 10 K lower as compared to $T_{g,\text{org}}$ by Eq. (2). $T_{g,\text{org}}$ at low RH is
370 predicted to be higher than $T_{g,\text{org}}$ at high RH, which results from a lower abundance of high molar
371 mass compounds observed at high RH. This trend is consistent with Kidd et al. (2014), who
372 showed that SOA material formed under dry conditions is more viscous than that formed under
373 wet conditions.

374 Figure 7(b) shows the predicted viscosity of toluene SOA as a function of RH, as
375 compared to the measured viscosity of toluene SOA formed in an oxidation flow reactor at 13%
376 RH and two mass loading ranges ($60 - 100 \mu\text{g m}^{-3}$ and $600 - 1000 \mu\text{g m}^{-3}$) (Song et al., 2016a).
377 Indirect viscosity measurements are also included in shaded boxes (Bateman et al., 2015; Li et



378 al., 2015). Lines with shaded areas are calculated viscosities using $T_{g,org}$ as described above. κ
379 was assumed to be 0.25 based on laboratory measurements (Lambe et al., 2011a; Hildebrandt
380 Ruiz et al., 2015). For achieving good fit, D was set to 13 and k_{GT} was assumed to be 3.0 (Dette
381 et al., 2014). Estimations with Eq. (1) match the measured viscosity values very well over the
382 entire RH range. Predictions with Eq. (2) overestimated the measurements by one or two orders
383 of magnitude at moderate RH between 30% and 50%, while they agreed with the measurements
384 derived at $RH \geq 60\%$ and at the dry conditions.

385 There are several possible reasons for the difference between the measurements and
386 predictions. First, the relative abundance of high molar mass compounds observed in HRMS
387 measurements may be overestimated, as high molar mass compounds tend to have higher (yet
388 generally unknown) ionization efficiencies compared to lower molar mass compounds. Second,
389 the nano-DESI-HRMS analysis of toluene SOA was limited to m/z range of 100 -1000 (Hinks et
390 al., 2017). It is possible that some SOA products with lower molar mass were present in particles
391 but not detected, which would lead to an overestimation of T_g . Third, the chemical composition
392 of toluene SOA are likely different between Hinks et al. (2017) and Song et al. (2016) because of
393 the differences in the experimental conditions. Specifically, toluene SOA was formed in a Teflon
394 chamber in Hinks et al., while Song et al. used an oxidation flow reactor to generate toluene
395 SOA. The O:C ratios are 0.71 at low RH and 0.63 at high RH based on nano-DESI-HRMS
396 measurements in Hinks et al. (2017), while it was 1.06 based on the aerosol mass spectrometry
397 (AMS) measurements in Song et al. (2016). Simultaneous measurements of viscosity and
398 chemical composition should be performed in future studies.

399

400 **3.3 Biomass Burning Particles**



401 To further explore the applicability of our viscosity prediction method using elemental
402 composition as measured by HRMS, we performed similar calculations for biomass burning
403 organic particles emitted from test facility burns of subalpine fir and lodgepole pine trees,
404 conducted as a part of the FIREX 2016 campaign (Selimovic et al., 2017). These samples were
405 analyzed by HRMS using two different ionization sources: electrospray ionization (ESI) and
406 atmospheric pressure photoionization (APPI). Mass spectra shown in Fig. 8(a) and (b) indicate
407 that a substantial number of compounds were detected by both methods (109 and 170
408 compounds for subalpine fir and lodgepole pine, respectively). However, pronounced
409 differences are also observed between the ESI and APPI spectra both in terms of the identity and
410 signal intensities of the detected compounds.

411 Glass transition temperatures for the assigned CH and CHO compounds were computed
412 using Eq. (2). Nitrogen and sulfur containing compounds (CHON and CHOS) are not yet
413 covered by Eq. (2) and were therefore excluded from the analysis. The exclusion is justified
414 because in the selected samples, CHON and CHOS compounds comprised less than 10% of the
415 detected ion intensity and <15% of the assigned compounds. T_g of organic mixtures ($T_{g,org}$) were
416 then calculated using the Gordon-Taylor approach with $k_{GT} = 1$, assuming that the relative
417 concentration of each compound is proportional to its MS signal intensity. The calculated $T_{g,org}$
418 values for the mixtures are specified in the legend of Figure 9. For both types of mixtures, the
419 calculated $T_{g,org}$ for the APPI MS data is lower than the value calculated based on the ESI MS
420 data with a difference of 32 K for subalpine fir and 11 K for the lodgepole pine. Figure 9 shows
421 the predicted viscosity as a function of RH, assuming $D = 10$, $\kappa = 0.10$ and $k_{GT} = 2.5$. The
422 difference in $T_{g,org}$ derived from ESI and APPI results in a variation of predicted viscosity at low
423 RH by up to five and two orders of magnitude for subalpine fir and lodgepole pine, respectively.



424 The difference in the calculated $T_{g,org}$ values is attributed to the chemical profile of the
425 species detected using different ionization techniques as shown in mass spectra in Fig. 8(a) and
426 (b). Van Krevelen diagrams in Fig. 8(c) and (d) illustrate these compositional differences
427 between chemical species detected by ESI and APPI. ESI is more efficient at detection of polar
428 compounds (Kiontke et al., 2016), which typically have higher O:C ratios and therefore would
429 result in higher predicted values of viscosity (Koop et al., 2011; Saukko et al., 2012). APPI
430 enables the detection of nonpolar compounds with lower O:C ratios, in particular polycyclic
431 aromatic hydrocarbons (PAHs), that have low ionization efficiencies when analyzed by ESI MS
432 (Raffaelli and Saba, 2003; Itoh et al., 2006). Due to the complementary nature of these ionization
433 methods, it is most likely that the actual glass transition temperature and viscosity of each type of
434 SOA are somewhere in between the values inferred from ESI and APPI data sets: ESI MS may
435 be viewed as providing the upper limit of viscosity, while APPI MS gives the lower limit. Our
436 results indicate that the use of complementary ionization techniques may help evaluate the
437 associated uncertainty for the prediction of viscosity values based on the elemental composition
438 as measured by HRMS.

439

440 **4 Conclusion**

441 We have developed a parameterization for calculation of the glass transition temperature
442 of individual SOA compounds with molar mass up to $\sim 1100 \text{ g mol}^{-1}$ using the number of carbon,
443 oxygen, and hydrogen atoms. Viscosity of SOA was estimated using the T_g -scaled Arrhenius plot
444 of viscosity versus T_g/T and the Gordon-Taylor approach to account for mixtures of SOA and
445 water. The fragility parameter D was compiled for organic compounds and we found that D
446 approaches a lower limit of ~ 10 (+/- 1.7) as the molar mass increases. The resulting viscosity



447 estimations agree well with measured viscosity of α -pinene and isoprene SOA, validating our
448 method. Using HRMS data, glass transition temperatures of individual components and viscosity
449 of toluene SOA were predicted, also resulting in a good agreement with measurements.
450 However, we note that the predicted viscosities were slightly higher than the measured values
451 suggesting that additional considerations may need to be taken into account. For example, the
452 ionization efficiency of both low and high molar mass compounds may have a pronounced effect
453 on the relative abundance of different classes of compounds in HRMS data. The viscosity
454 prediction method was also applied to biomass burning particles, whose elemental composition
455 was measured using HRMS with two different ionization techniques. Substantial differences in
456 viscosity estimations were obtained using ESI and APPI mass spectra.

457 Figure 10 summarizes the predicted range of viscosity of α -pinene SOA, isoprene SOA,
458 toluene SOA, and biomass burning particles. Isoprene SOA has lower viscosity, reflecting lower
459 glass transition temperature due to relatively low molar mass of isoprene oxidation products. α -
460 pinene and toluene SOA have much higher viscosity with a different shape of the RH
461 dependence due to differences in glass transition temperatures and hygroscopicity. Biomass
462 burning particles have moderate viscosity between the two extreme cases. Currently, both
463 predictions and measurements are subject to large uncertainties and variations. Complementary
464 measurements of viscosity and chemical composition employing different ionization techniques
465 are desired to further constrain RH-dependent viscosity in future studies. Current T_g
466 parameterizations do not consider functionality or molecular structure explicitly and further
467 measurements of T_g and viscosity of SOA would allow us to refine the method presented in this
468 study. Nevertheless, current results offer a promising starting point and such simple
469 parameterizations are practical for predicting viscosity of particles as measured by HRMS. The



470 developed viscosity prediction method should also be useful in recent efforts of simulating the
 471 distribution of SOA phase state and related properties in regional or global air quality models
 472 (e.g., Maclean et al., 2017; Shiraiwa et al., 2017).

473

474 **Appendix A: Conversion of fragility steepness index (m) to fragility (D)**

475 Fragility steepness index (m) is defined as:

$$476 \quad m = \lim_{T \rightarrow T_g} \frac{d \log \eta}{d(T_g/T)} \quad (\text{A1})$$

477 Combining Eq. (A1) with Eq. (4) gives:

$$478 \quad m = \lim_{T \rightarrow T_g} \frac{d}{d(T_g/T)} \left(-5 + 0.434 \frac{T_0 D}{T - T_0} \right) \quad (\text{A2})$$

479 Considering that $\eta = 10^{12}$ Pa s at $T = T_g$ (Angell, 1991), and by defining $\Delta x = 1 - T_g/T$, and a
 480 combination with Eq. (5) leads to:

$$\begin{aligned}
 m &= \lim_{\Delta x \rightarrow 0} \frac{1}{\Delta x} \left(12 - \left(-5 + 0.434 \frac{\frac{39.17 T_g}{D + 39.17} D}{\frac{T_g}{1 - \Delta x} - \frac{39.17 T_g}{D + 39.17}} \right) \right) \\
 &= \lim_{\Delta x \rightarrow 0} \frac{1}{\Delta x} \left(17 - 0.434 \frac{39.17 T_g D (1 - \Delta x)}{D T_g + 39.17 T_g \Delta x} \right) \\
 &= \lim_{\Delta x \rightarrow 0} \frac{(665.89 + 17D)}{(D + 39.17 \Delta x)} \\
 &= \frac{665.89 + 17D}{D} \quad (\text{A3})
 \end{aligned}$$

482 Note that Eq. (A3) is derived assuming the high temperature limit of viscosity η_∞ is equal to 10^{-5}
 483 Pa s (Angell, 1991) in the VTF equation (Eq. 3). Similar equations for the relation between m
 484 and D were given by previous studies using different η_∞ and units (Angell et al., 1994; Angell,
 485 2002; Bones et al., 2012) and applying those gave very similar results in our study.

486



487 **Acknowledgements.**

488 This work was funded by the National Science Foundation (AGS-1654104) and the Department
489 of Energy (DE-SC0018349). The Purdue group and S. N. acknowledge additional support by the
490 U.S. Department of Commerce, National Oceanic and Atmospheric Administration through
491 Climate Program Office's AC4 program, awards NA16OAR4310101 and NA16OAR4310102.
492 We thank Ulrich Pöschl and Thomas Koop for stimulating discussions.

493

494 **References.**

- 495 Abbatt, J. P. D., Lee, A. K. Y., and Thornton, J. A.: Quantifying trace gas uptake to tropospheric
496 aerosol: recent advances and remaining challenges, *Chem. Soc. Rev.*, 41, 6555-6581, 2012.
- 497 Abramson, E., Imre, D., Beranek, J., Wilson, J., and Zelenyuk, A.: Experimental determination
498 of chemical diffusion within secondary organic aerosol particles, *Physical Chemistry Chemical
499 Physics*, 15, 2983-2991, 2013.
- 500 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., et al.: O/C and OM/OC Ratios of Primary, Secondary,
501 and Ambient Organic Aerosols with High-Resolution Time-of-Flight Aerosol Mass
502 Spectrometry, *Environ. Sci. Technol.*, 42, 4478-4485, 2008.
- 503 Angell, C.: Relaxation in liquids, polymers and plastic crystals—strong/fragile patterns and
504 problems, *Journal of Non-Crystalline Solids*, 131, 13-31, 1991.
- 505 Angell, C. A., Bressel, R. D., Green, J. L., Kanno, H., Oguni, M., and Sare, E. J.: Liquid fragility
506 and the glass transition in water and aqueous solutions, *Journal of Food Engineering*, 22, 115-
507 142, 1994.
- 508 Angell, C. A.: Formation of glasses from liquids and biopolymers, *Science*, 267, 1924-1935,
509 1995.
- 510 Angell, C. A.: Entropy and fragility in supercooling liquids, National Institute of Standards and
511 Technology, *Journal of Research*, 102, 171-185, 1997.
- 512 Angell, C. A.: Liquid Fragility and the Glass Transition in Water and Aqueous Solutions, *Chem.
513 Rev.*, 102, 2627-2650, 2002.
- 514 Arangio, A. M., Slade, J. H., Berkemeier, T., Pöschl, U., Knopf, D. A., and Shiraiwa, M.:
515 Multiphase Chemical Kinetics of OH Radical Uptake by Molecular Organic Markers of Biomass
516 Burning Aerosols: Humidity and Temperature Dependence, *Surface Reaction and Bulk
517 Diffusion*, *J. Phys. Chem. A*, 119, 4533-4544, 2015.



- 518 Bateman, A. P., Laskin, J., Laskin, A., and Nizkorodov, S. A.: Applications of High-Resolution
519 Electrospray Ionization Mass Spectrometry to Measurements of Average Oxygen to Carbon
520 Ratios in Secondary Organic Aerosols, *Environ. Sci. Technol.*, 46, 8315-8324, 2012.
- 521 Bateman, A. P., Bertram, A. K., and Martin, S. T.: Hygroscopic Influence on the Semisolid-to-
522 Liquid Transition of Secondary Organic Materials, *J. Phys. Chem. A*, 119, 4386-4395, 2015.
- 523 Bateman, A. P., Gong, Z., Liu, P., et al.: Sub-micrometre particulate matter is primarily in liquid
524 form over Amazon rainforest, *Nat. Geosci.*, 9, 34-37, 2016.
- 525 Baustian, K. J., Wise, M. E., Jensen, E. J., Schill, G. P., Freedman, M. A., and Tolbert, M. A.:
526 State transformations and ice nucleation in amorphous (semi-)solid organic aerosol, *Atmos.*
527 *Chem. Phys.*, 13, 5615-5628, 2013.
- 528 Berkemeier, T., Shiraiwa, M., Pöschl, U., and Koop, T.: Competition between water uptake and
529 ice nucleation by glassy organic aerosol particles, *Atmos. Chem. Phys.*, 14, 12513-12531, 2014.
- 530 Berkemeier, T., Steimer, S., Krieger, U. K., Peter, T., Poschl, U., Ammann, M., and Shiraiwa,
531 M.: Ozone uptake on glassy, semi-solid and liquid organic matter and the role of reactive oxygen
532 intermediates in atmospheric aerosol chemistry, *Phys. Chem. Chem. Phys.*, 18, 12662-12674,
533 2016.
- 534 Boehmer, R., Ngai, K. L., Angell, C. A., and Plazek, D. J.: Nonexponential relaxations in strong
535 and fragile glass formers, *J. Chem. Phys.*, 99, 4201-4209, 1993.
- 536 Bones, D. L., Reid, J. P., Lienhard, D. M., and Krieger, U. K.: Comparing the mechanism of
537 water condensation and evaporation in glassy aerosol, *Proc. Natl. Acad. Sci. U.S.A.*, 109, 11613-
538 11618, 2012.
- 539 Booth, A. M., Murphy, B., Riipinen, I., Percival, C. J., and Topping, D. O.: Connecting Bulk
540 Viscosity Measurements to Kinetic Limitations on Attaining Equilibrium for a Model Aerosol
541 Composition, *Environ. Sci. Technol.*, 48, 9298-9305, 2014.
- 542 Cappa, C. D., and Wilson, K. R.: Evolution of organic aerosol mass spectra upon heating:
543 implications for OA phase and partitioning behavior, *Atmos. Chem. Phys.*, 11, 1895-1911, 2011.
- 544 Cappa, C. D., and Wilson, K. R.: Multi-generation gas-phase oxidation, equilibrium partitioning,
545 and the formation and evolution of secondary organic aerosol, *Atmos. Chem. Phys.*, 12, 9505-
546 9528, 2012.
- 547 Chen, Q., Liu, Y., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-Phase Chemistry of
548 Secondary Organic Material: Modeled Compared to Measured O:C and H:C Elemental Ratios
549 Provide Constraints, *Environ. Sci. Technol.*, 45, 4763-4770, 2011.
- 550 Davies, J. F., and Wilson, K. R.: Nanoscale interfacial gradients formed by the reactive uptake of
551 OH radicals onto viscous aerosol surfaces, *Chem. Sci.*, 6, 7020-7027, 2015.



- 552 Dette, H. P., Qi, M., Schröder, D. C., Godt, A., and Koop, T.: Glass-forming properties of 3-
553 Methylbutane-1,2,3-tricarboxylic acid and its mixtures with water and pinonic acid, *J. Phys.*
554 *Chem. A*, 118, 7024-7033, 2014.
- 555 Dette, H. P., and Koop, T.: Glass Formation Processes in Mixed Inorganic/Organic Aerosol
556 Particles, *J. Phys. Chem. A*, 119, 4552-4561, 2015.
- 557 Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning,
558 dilution, and chemical aging of semivolatile organics, *Environ. Sci. Technol.*, 40, 2635-2643,
559 2006.
- 560 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility
561 basis set: 1. organic-aerosol mixing thermodynamics, *Atmos. Chem. Phys.*, 11, 3303-3318, 2011.
- 562 Fox Jr, T. G., and Flory, P. J.: Second - order transition temperatures and related properties of
563 polystyrene. I. Influence of molecular weight, *J. Appl. Phys.*, 21, 581-591, 1950.
- 564 Gao, S., Ng, N. L., Keywood, M., et al.: Particle phase acidity and oligomer formation in
565 secondary organic aerosol, *Environ. Sci. Technol.*, 38, 6582-6589, 2004.
- 566 Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the earth's
567 atmosphere, *Environ. Sci. Technol.*, 41, 1514-1521, 2007.
- 568 Grayson, J. W., Zhang, Y., Mutzel, A., Renbaum-Wolff, L., Böge, O., Kamal, S., Herrmann, H.,
569 Martin, S. T., and Bertram, A. K.: Effect of varying experimental conditions on the viscosity of
570 α -pinene derived secondary organic material, *Atmos. Chem. Phys.*, 16, 6027-6040, 2016.
- 571 Grayson, J. W., Evoy, E., Song, M., et al.: The effect of hydroxyl functional groups and molar
572 mass on the viscosity of non-crystalline organic and organic-water particles, *Atmos. Chem.*
573 *Phys.*, 17, 8509-8524, 2017.
- 574 Gunthe, S. S., King, S. M., Rose, D., et al.: Cloud condensation nuclei in pristine tropical
575 rainforest air of Amazonia: size-resolved measurements and modeling of atmospheric aerosol
576 composition and CCN activity, *Atmos. Chem. Phys.*, 9, 7551-7575, 2009.
- 577 Hancock, B. C., and Zografi, G.: The relationship between the glass transition temperature and
578 the water content of amorphous pharmaceutical solids, *Pharm. Res.*, 11, 471-477, 1994.
- 579 Hildebrandt Ruiz, L., Paciga, A. L., Cerully, K. M., Nenes, A., Donahue, N. M., and Pandis, S.
580 N.: Formation and aging of secondary organic aerosol from toluene: changes in chemical
581 composition, volatility, and hygroscopicity, *Atmos. Chem. Phys.*, 15, 8301-8313, 2015.
- 582 Hinks, M. L., Brady, M. V., Lignell, H., et al.: Effect of viscosity on photodegradation rates in
583 complex secondary organic aerosol materials, *Phys. Chem. Chem. Phys.*, 18, 8785-8793, 2016.
- 584 Hinks, M. L., Montoya-Aguilera, J., Ellison, L., Lin, P., Laskin, A., Laskin, J., Shiraiwa, M.,
585 Dabdub, D., and Nizkorodov, S. A.: Effect of Relative Humidity on the Composition of



- 586 Secondary Organic Aerosol from Oxidation of Toluene, Atmos. Chem. Phys. Discuss., 2017, 1-
587 16, 2017.
- 588 Hosny, N. A., Fitzgerald, C., Vysniauskas, A., et al.: Direct imaging of changes in aerosol
589 particle viscosity upon hydration and chemical aging, Chem. Sci., 7, 1357-1367, 2016.
- 590 Huang, D., and McKenna, G. B.: New insights into the fragility dilemma in liquids, J. Chem.
591 Phys., 114, 5621-5630, 2001.
- 592 Ignatius, K., Kristensen, T. B., Järvinen, E., et al.: Heterogeneous ice nucleation of viscous
593 secondary organic aerosol produced from ozonolysis of α -pinene, Atmos. Chem. Phys., 16,
594 6495-6509, 2016.
- 595 Itoh, N., Aoyagi, Y., and Yarita, T.: Optimization of the dopant for the trace determination of
596 polycyclic aromatic hydrocarbons by liquid chromatography/dopant-assisted atmospheric-
597 pressure photoionization/mass spectrometry, J. Chromatogr. A, 1131, 285-288, 2006.
- 598 Jain, S., and Petrucci, G. A.: A New Method to Measure Aerosol Particle Bounce Using a
599 Cascade Electrical Low Pressure Impactor, Aerosol Sci. Technol., 49, 390-399, 2015.
- 600 Jathar, S. H., Cappa, C. D., Wexler, A. S., Seinfeld, J. H., and Kleeman, M. J.: Multi-
601 generational oxidation model to simulate secondary organic aerosol in a 3-D air quality model,
602 Geosci. Model Dev., 8, 2553-2567, 2015.
- 603 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., et al.: Evolution of organic aerosols in the
604 atmosphere, Science, 326, 1525-1529, 2009.
- 605 Kidd, C., Perraud, V., Wingen, L. M., and Finlayson-Pitts, B. J.: Integrating phase and
606 composition of secondary organic aerosol from the ozonolysis of alpha-pinene, Proc. Natl. Acad.
607 Sci. U.S.A., 111, 7552-7557, 2014.
- 608 Kiontke, A., Oliveira-Birkmeier, A., Opitz, A., and Birkemeyer, C.: Electrospray ionization
609 efficiency is dependent on different molecular descriptors with respect to solvent pH and
610 instrumental configuration, PLoS One, 11, e0167502/0167501-e0167502/0167516, 2016.
- 611 Koop, T., Bookhold, J., Shiraiwa, M., and Pöschl, U.: Glass transition and phase state of organic
612 compounds: dependency on molecular properties and implications for secondary organic
613 aerosols in the atmosphere, Phys. Chem. Chem. Phys., 13, 19238-19255, 2011.
- 614 Kuwata, M., and Martin, S. T.: Phase of atmospheric secondary organic material affects its
615 reactivity, Proc. Natl. Acad. Sci. U.S.A., 109, 17354-17359, 2012.
- 616 Kuwata, M., Zorn, S. R., and Martin, S. T.: Using elemental ratios to predict the density of
617 organic material composed of carbon, hydrogen, and oxygen, Environ. Sci. Technol., 46, 787-
618 794, 2012.



- 619 Lambe, A. T., Ahern, A. T., Williams, L. R., et al.: Characterization of aerosol photooxidation
620 flow reactors: heterogeneous oxidation, secondary organic aerosol formation and cloud
621 condensation nuclei activity measurements, *Atmos. Meas. Tech.*, 4, 445-461, 2011a.
- 622 Lambe, A. T., Onasch, T. B., Massoli, P., et al.: Laboratory studies of the chemical composition
623 and cloud condensation nuclei (CCN) activity of secondary organic aerosol (SOA) and oxidized
624 primary organic aerosol (OPOA), *Atmos. Chem. Phys.*, 11, 8913-8928, 2011b.
- 625 Li, Y., Pöschl, U., and Shiraiwa, M.: Molecular corridors and parameterizations of volatility in
626 the chemical evolution of organic aerosols, *Atmos. Chem. Phys.*, 16, 3327-3344, 2016.
- 627 Li, Y. J., Liu, P., Gong, Z., Wang, Y., Bateman, A. P., Bergoend, C., Bertram, A. K., and Martin,
628 S. T.: Chemical Reactivity and Liquid/Nonliquid States of Secondary Organic Material, *Environ.*
629 *Sci. Technol.*, 49, 13264-13274, 2015.
- 630 Lienhard, D. M., Huisman, A. J., Krieger, U. K., et al.: Viscous organic aerosol particles in the
631 upper troposphere: diffusivity-controlled water uptake and ice nucleation?, *Atmos. Chem. Phys.*,
632 15, 13599-13613, 2015.
- 633 Lignell, H., Hinks, M. L., and Nizkorodov, S. A.: Exploring matrix effects on photochemistry of
634 organic aerosols, *Proc. Natl. Acad. Sci. U.S.A.*, 111, 13780-13785, 2014.
- 635 Liu, P., Li, Y. J., Wang, Y., Gilles, M. K., Zaveri, R. A., Bertram, A. K., and Martin, S. T.:
636 Lability of secondary organic particulate matter, *Proc. Natl. Acad. Sci. U.S.A.*, 113, 12643-
637 12648, 2016.
- 638 Maclean, A. M., Butenhoff, C. L., Grayson, J. W., Barsanti, K., Jimenez, J. L., and Bertram, A.
639 K.: Mixing times of organic molecules within secondary organic aerosol particles: a global
640 planetary boundary layer perspective, *Atmos. Chem. Phys.*, 17, 13037-13048, 2017.
- 641 Marshall, F. H., Miles, R. E. H., Song, Y.-C., Ohm, P. B., Power, R. M., Reid, J. P., and Dutcher,
642 C. S.: Diffusion and reactivity in ultraviscous aerosol and the correlation with particle viscosity,
643 *Chem. Sci.*, 7, 1298-1308, 2016.
- 644 Matsushima, S., Takano, A., Takahashi, Y., and Matsushita, Y.: Precise synthesis of a series of
645 poly(4-n-alkylstyrene)s and their glass transition temperatures, *Journal of Polymer Science Part*
646 *B: Polymer Physics*, 55, 757-763, 2017.
- 647 Mikhailov, E., Vlasenko, S., Martin, S. T., Koop, T., and Pöschl, U.: Amorphous and crystalline
648 aerosol particles interacting with water vapor: conceptual framework and experimental evidence
649 for restructuring, phase transitions and kinetic limitations, *Atmos. Chem. Phys.*, 9, 9491-9522,
650 2009.
- 651 Montserrat, S., and Colomer, P.: The effect of the molecular weight on the glass transition
652 temperature in amorphous poly(ethylene terephthalate), *Polymer Bulletin*, 12, 173-180, 1984.
- 653 Murray, B. J., Wilson, T. W., Dobbie, S., et al.: Heterogeneous nucleation of ice particles on
654 glassy aerosols under cirrus conditions, *Nat. Geosci.*, 3, 233-237, 2010.



- 655 Nakanishi, M., and Nozaki, R.: Systematic study of the glass transition in polyhydric alcohols,
656 Physical Review E, 83, 051503, 2011.
- 657 Nguyen, T. B., Nizkorodov, S. A., Laskin, A., and Laskin, J.: An approach toward quantification
658 of organic compounds in complex environmental samples using high-resolution electrospray
659 ionization mass spectrometry, Anal. Methods, 5, 72-80, 2013.
- 660 Nizkorodov, S. A., Laskin, J., and Laskin, A.: Molecular chemistry of organic aerosols through
661 the application of high resolution mass spectrometry, Phys. Chem. Chem. Phys., 13, 3612-3629,
662 2011.
- 663 Nozière, B., Kalberer, M., Claeys, M., et al.: The Molecular Identification of Organic
664 Compounds in the Atmosphere: State of the Art and Challenges, Chem. Rev., 115, 3919–3983,
665 2015.
- 666 O’Connell, P. A., and McKenna, G. B.: Arrhenius-type temperature dependence of the segmental
667 relaxation below T_g , J. Chem. Phys., 110, 11054-11060, 1999.
- 668 Onder, K., Peters, R. H., and Spark, L. C.: Melting and transition phenomena in some polyester-
669 urethanes, Polymer, 13, 133-139, 1972.
- 670 Pajunoja, A., Malila, J., Hao, L., Joutsensaari, J., Lehtinen, K. E. J., and Virtanen, A.: Estimating
671 the viscosity range of SOA particles based on their coalescence time, Aerosol Sci. Technol., 48,
672 i-iv, 2014.
- 673 Pajunoja, A., Lambe, A. T., Hakala, J., et al.: Adsorptive uptake of water by semisolid secondary
674 organic aerosols, Geophys. Res. Lett., 42, 3063-3068, 2015.
- 675 Papadopoulos, P., Floudas, G., Chi, C., and Wegner, G.: Molecular dynamics of oligofluorenes:
676 A dielectric spectroscopy investigation, J. Chem. Phys., 120, 2368-2374, 2004.
- 677 Perraud, V., Bruns, E. A., Ezell, M. J., et al.: Nonequilibrium atmospheric secondary organic
678 aerosol formation and growth, Proc. Natl. Acad. Sci. U.S.A., 109, 2836-2841, 2012.
- 679 Petters, M. D., and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth
680 and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961-1971, 2007.
- 681 Petters, S. S., Pagonis, D., Claflin, M. S., Levin, E. J. T., Petters, M. D., Ziemann, P. J., and
682 Kreidenweis, S. M.: Hygroscopicity of Organic Compounds as a Function of Carbon Chain
683 Length and Carboxyl, Hydroperoxy, and Carbonyl Functional Groups, J. Phys. Chem. A, 121,
684 5164-5174, 2017.
- 685 Polymer handbook, t. e.: J. Brandrup (Editor), Edmund H. Immergut (Editor), E. A. Grulke
686 (Editor), John Wiley & Sons, Inc., ISBN 0-471-16628-6, 1999.
- 687 Price, H. C., Murray, B. J., Mattsson, J., O’Sullivan, D., Wilson, T. W., Baustian, K. J., and
688 Benning, L. G.: Quantifying water diffusion in high-viscosity and glassy aqueous solutions using
689 a Raman isotope tracer method, Atmos. Chem. Phys., 14, 3817-3830, 2014.



- 690 Putman, A. L., Offenberg, J. H., Fisseha, R., Kundu, S., Rahn, T. A., and Mazzoleni, L. R.:
691 Ultrahigh-resolution FT-ICR mass spectrometry characterization of α -pinene ozonolysis SOA,
692 Atmos. Environ., 46, 164-172, 2012.
- 693 Raffaelli, A., and Saba, A.: Atmospheric pressure photoionization mass spectrometry, Mass
694 Spectrom. Rev., 22, 318-331, 2003.
- 695 Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, K., Sellier, M., Murray, B. J.,
696 Schilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of α -pinene secondary organic
697 material and implications for particle growth and reactivity, Proc. Natl. Acad. Sci. U.S.A., 110,
698 8014-8019, 2013.
- 699 Roach, P. J., Laskin, J., and Laskin, A.: Nanospray desorption electrospray ionization: an
700 ambient method for liquid-extraction surface sampling in mass spectrometry, Analyst
701 (Cambridge, U. K.), 135, 2233-2236, 2010a.
- 702 Roach, P. J., Laskin, J., and Laskin, A.: Molecular Characterization of Organic Aerosols Using
703 Nanospray-Desorption/Electrospray Ionization-Mass Spectrometry, Anal. Chem. (Washington,
704 DC, U. S.), 82, 7979-7986, 2010b.
- 705 Roos, Y.: Melting and glass transitions of low molecular weight carbohydrates, Carbohydr. Res.,
706 238, 39-48, 1993.
- 707 Rothfuss, N. E., and Petters, M. D.: Influence of Functional Groups on the Viscosity of Organic
708 Aerosol, Environ. Sci. Technol., 51, 271-279, 2017.
- 709 Saukko, E., Lambe, A. T., Massoli, P., et al.: Humidity-dependent phase state of SOA particles
710 from biogenic and anthropogenic precursors, Atmos. Chem. Phys., 12, 7517-7529, 2012.
- 711 Schill, G. P., and Tolbert, M. A.: Heterogeneous ice nucleation on phase-separated organic-
712 sulfate particles: effect of liquid vs. glassy coatings, Atmos. Chem. Phys., 13, 4681-4695, 2013.
- 713 Schill, G. P., De Haan, D. O., and Tolbert, M. A.: Heterogeneous Ice Nucleation on Simulated
714 Secondary Organic Aerosol, Environ. Sci. Technol., 48, 1675-1682, 2014.
- 715 Selimovic, V., Yokelson, R. J., Warneke, C., Roberts, J. M., de Gouw, J., Reardon, J., and
716 Griffith, D. W. T.: Aerosol optical properties and trace gas emissions by PAX and OP-FTIR for
717 laboratory-simulated western US wildfires during FIREX, Atmos. Chem. Phys. Discuss., 2017,
718 1-34, 2017.
- 719 Shiraiwa, M., Ammann, M., Koop, T., and Pöschl, U.: Gas uptake and chemical aging of
720 semisolid organic aerosol particles, Proc. Natl. Acad. Sci. U.S.A., 108, 11003-11008, 2011.
- 721 Shiraiwa, M., and Seinfeld, J. H.: Equilibration timescale of atmospheric secondary organic
722 aerosol partitioning, Geophys. Res. Lett., 39, L24801, 2012.



- 723 Shiraiwa, M., Zuend, A., Bertram, A. K., and Seinfeld, J. H.: Gas-particle partitioning of
724 atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology, *Phys.*
725 *Chem. Chem. Phys.*, 15, 11441-11453, 2013.
- 726 Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K. A., Seinfeld, J. H., and Pöschl, U.:
727 Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary
728 organic aerosol, *Atmos. Chem. Phys.*, 14, 8323-8341, 2014.
- 729 Shiraiwa, M., Li, Y., Tsimpidi, A. P., Karydis, V. A., Berkemeier, T., Pandis, S. N., Lelieveld, J.,
730 Koop, T., and Pöschl, U.: Global distribution of particle phase state in atmospheric secondary
731 organic aerosols, *Nat. Commun.*, 8, 15002, 2017.
- 732 Shrivastava, M., Cappa, C. D., Fan, J., et al.: Recent advances in understanding secondary
733 organic aerosol: Implications for global climate forcing, *Rev. Geophys.*, 55, 509-559, 2017.
- 734 Slade, J. H., and Knopf, D. A.: Multiphase OH Oxidation Kinetics of Organic Aerosol: The Role
735 of Particle Phase State and Relative Humidity, *Geophys. Res. Lett.*, 2014GL060582, 2014.
- 736 Song, M., Liu, P. F., Hanna, S. J., Li, Y. J., Martin, S. T., and Bertram, A. K.: Relative humidity-
737 dependent viscosities of isoprene-derived secondary organic material and atmospheric
738 implications for isoprene-dominant forests, *Atmos. Chem. Phys.*, 15, 5145-5159, 2015.
- 739 Song, M., Liu, P. F., Hanna, S. J., Zaveri, R. A., Potter, K., You, Y., Martin, S. T., and Bertram,
740 A. K.: Relative humidity-dependent viscosity of secondary organic material from toluene photo-
741 oxidation and possible implications for organic particulate matter over megacities, *Atmos. Chem.*
742 *Phys.*, 16, 8817-8830, 2016a.
- 743 Song, Y. C., Haddrell, A. E., Bzdek, B. R., Reid, J. P., Bannan, T., Topping, D. O., Percival, C.,
744 and Cai, C.: Measurements and Predictions of Binary Component Aerosol Particle Viscosity, *J.*
745 *Phys. Chem. A*, 120, 8123-8137, 2016b.
- 746 Tolocka, M. P., Jang, M., Ginter, J. M., Cox, F. J., Kamens, R. M., and Johnston, M. V.:
747 Formation of oligomers in secondary organic aerosol, *Environ. Sci. Technol.*, 38, 1428-1434,
748 2004.
- 749 US-EPA: Estimation programs interface suite for microsoft windows, 2012.
- 750 Vaden, T. D., Imre, D., Beranek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation kinetics and
751 phase of laboratory and ambient secondary organic aerosol, *Proc. Natl. Acad. Sci. U.S.A.*, 108,
752 2190-2195, 2011.
- 753 van der Sman, R. G. M.: Predictions of Glass Transition Temperature for Hydrogen Bonding
754 Biomaterials, *J. Phys. Chem. B*, 117, 16303-16313, 2013.
- 755 Virtanen, A., Joutsensaari, J., Koop, T., et al.: An amorphous solid state of biogenic secondary
756 organic aerosol particles, *Nature*, 467, 824-827, 2010.



- 757 Wang, B., O'Brien, R. E., Kelly, S. T., Shilling, J. E., Moffet, R. C., Gilles, M. K., and Laskin,
758 A.: Reactivity of Liquid and Semisolid Secondary Organic Carbon with Chloride and Nitrate in
759 Atmospheric Aerosols, *J. Phys. Chem. A*, 119, 4498-4508, 2015.
- 760 Wang, B. B., Lambe, A. T., Massoli, P., Onasch, T. B., Davidovits, P., Worsnop, D. R., and
761 Knopf, D. A.: The deposition ice nucleation and immersion freezing potential of amorphous
762 secondary organic aerosol: Pathways for ice and mixed-phase cloud formation, *J. Geophys. Res.-*
763 *Atmos.*, 117, D16209, 2012.
- 764 Williams, M. L., Landel, R. F., and Ferry, J. D.: The temperature dependence of relaxation
765 mechanisms in amorphous polymers and other glass-forming liquids, *J. Am. Chem. Soc.*, 77,
766 3701-3707, 1955.
- 767 Yli-Juuti, T., Pajunoja, A., Tikkanen, O.-P., et al.: Factors controlling the evaporation of
768 secondary organic aerosol from α -pinene ozonolysis, *Geophys. Res. Lett.*, 44, 2562-2570, 2017.
- 769 Zaveri, R. A., Easter, R. C., Shilling, J. E., and Seinfeld, J. H.: Modeling kinetic partitioning of
770 secondary organic aerosol and size distribution dynamics: representing effects of volatility, phase
771 state, and particle-phase reaction, *Atmos. Chem. Phys.*, 14, 5153-5181, 2014.
- 772 Zhang, Y., Sanchez, M. S., Douet, C., et al.: Changing shapes and implied viscosities of
773 suspended submicron particles, *Atmos. Chem. Phys.*, 15, 7819-7829, 2015.
- 774 Zhou, S., Shiraiwa, M., McWhinney, R., Pöschl, U., and Abbatt, J. P. D.: Kinetic limitations in
775 gas-particle reactions arising from slow diffusion in secondary organic aerosol, *Faraday Discuss.*,
776 165, 391-406, 2013.
- 777 Zobrist, B., Marcolli, C., Pedernera, D. A., and Koop, T.: Do atmospheric aerosols form
778 glasses?, *Atmos. Chem. Phys.*, 8, 5221-5244, 2008.
- 779 Zobrist, B., Soonsin, V., Luo, B. P., Krieger, U. K., Marcolli, C., Peter, T., and Koop, T.: Ultra-
780 slow water diffusion in aqueous sucrose glasses, *Phys. Chem. Chem. Phys.*, 13, 3514-3526,
781 2011.
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783



784 **Table 1.** Composition classes and the n_C^0 and b values (K) for glass transition temperature
 785 parameterizations obtained by least-squares optimization using the measurements compiled in
 786 Koop et al., (2011), Dette et al., (2014 and Rothfuss and Petters (2017).

Classes	n_C^0	b_C	b_H	b_{CH}	b_O	b_{CO}
CH	1.96 (±1.81)	61.99 (±53.65)	-113.33 (±44.47)	28.74 (±20.86)		
CHO	12.13 (±2.66)	10.95 (±13.60)	-41.82 (±14.78)	21.61 (±5.30)	118.96 (±9.72)	-24.38 (±4.21)

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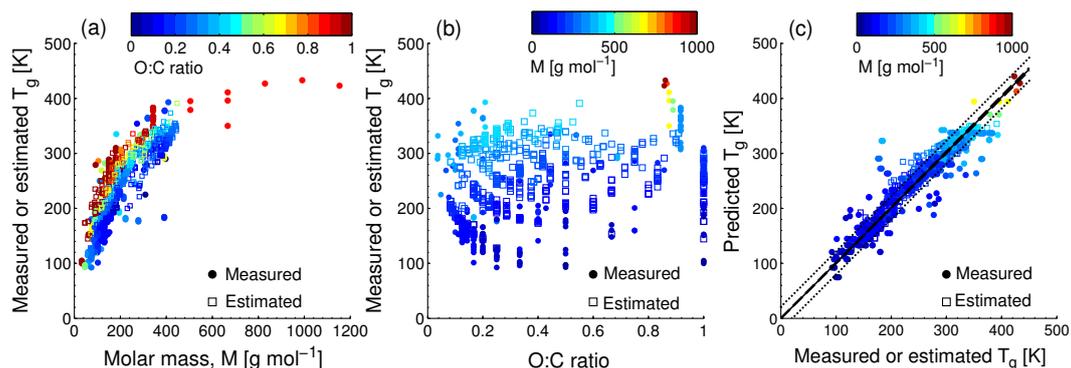
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790 **Table 2.** Glass transition temperatures calculated using Eq. (1) and (2) for toluene SOA mixtures
 791 at low relative humidity (low RH < 2%) and high relative humidity (high RH = 75%) conditions.

$T_{g,org}$ (K)	low RH	high RH
Equation (1)*	299	295
Equation (2)	313	303

792 * Compounds with $M > 450 \text{ g mol}^{-1}$ were excluded from the analysis.



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795 **Figure 1.** Characteristic relationships between molecular properties and the glass transition796 temperature (T_g) of organic compounds. (a) T_g of organic compounds as measured (circles) and

797 estimated with the Boyer-Kauzmann rule (squares) plotted against molar mass. The markers are

798 color-coded by atomic O:C ratio. (b) Measured (circles) and estimated (squares) T_g of organic

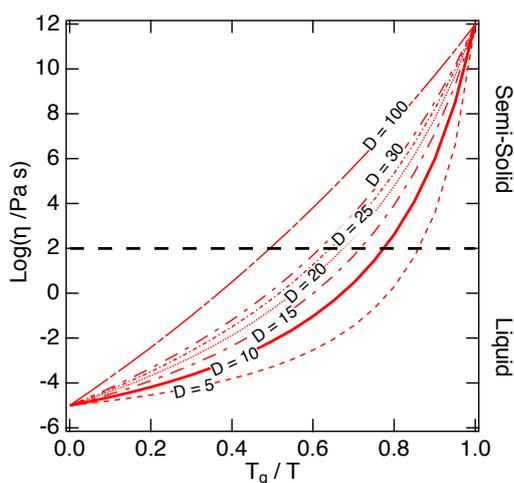
799 compounds plotted against O:C ratio. The markers are color-coded by molar mass. (c) Predicted

800 T_g for CHO compounds using a parameterization (Eq. 2) developed in this study compared to801 measured (circles) and estimated T_g by the Boyer-Kauzmann rule (squares). The solid line shows

802 1:1 line and the dashed and dotted lines show 68% confidence and prediction bands,

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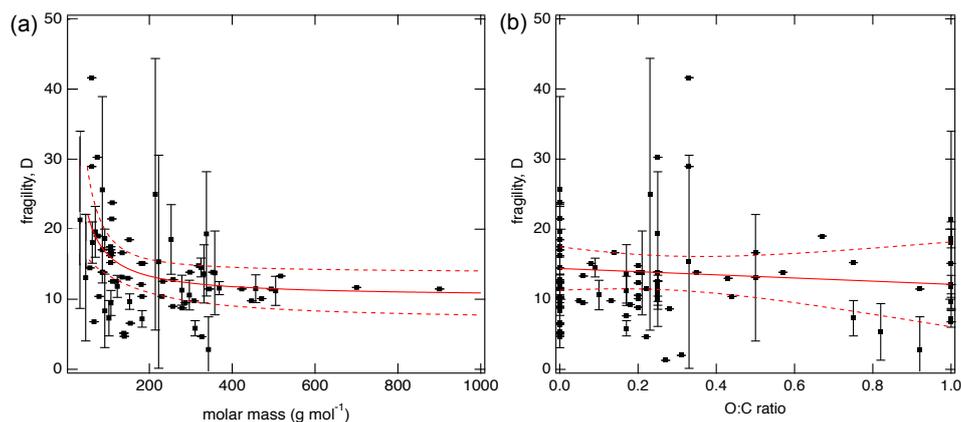
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806 **Figure 2.** The Angell plot of viscosity (η) vs. T_g/T . The lines represent different fragility807 parameter (D) values in the range of 5 - 100, with $D = 10$ (the solid line) used as a base case for

808 this study. A large fragility parameter value is associated with a strong glass former, while

809 fragile materials are associated with lower values. The black dashed line at viscosity of 10^2 Pa s

810 indicates the approximate threshold between liquid and semi-solid states.



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812 **Figure 3.** Fragility parameter of organic compounds (D) plotted against (a) molar mass and (b)

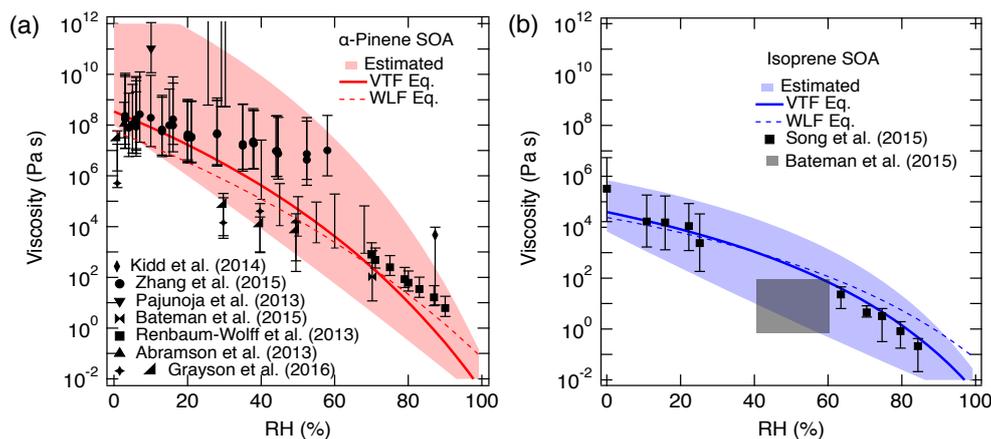
813 atomic O:C ratio. Error bars are standard deviations. The solid red lines represent the fitted

814 curves with fitted equations for (a) $D = 602.6/M + 10.3$ and (b) $D = 14.4 - 2.3(\text{O:C})$ respectively.

815 Dashed red lines indicate the 95% confidence band.

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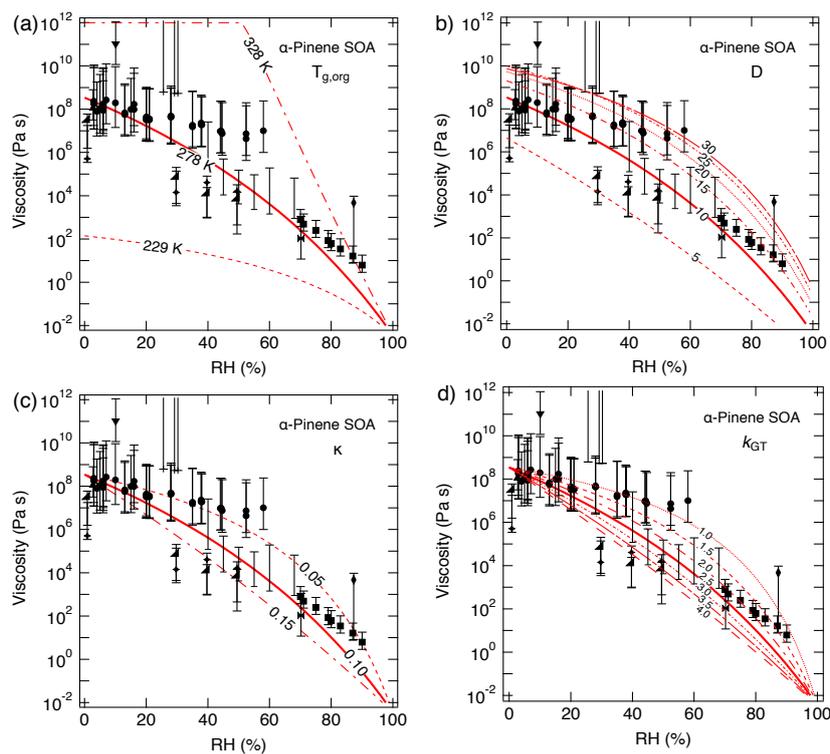
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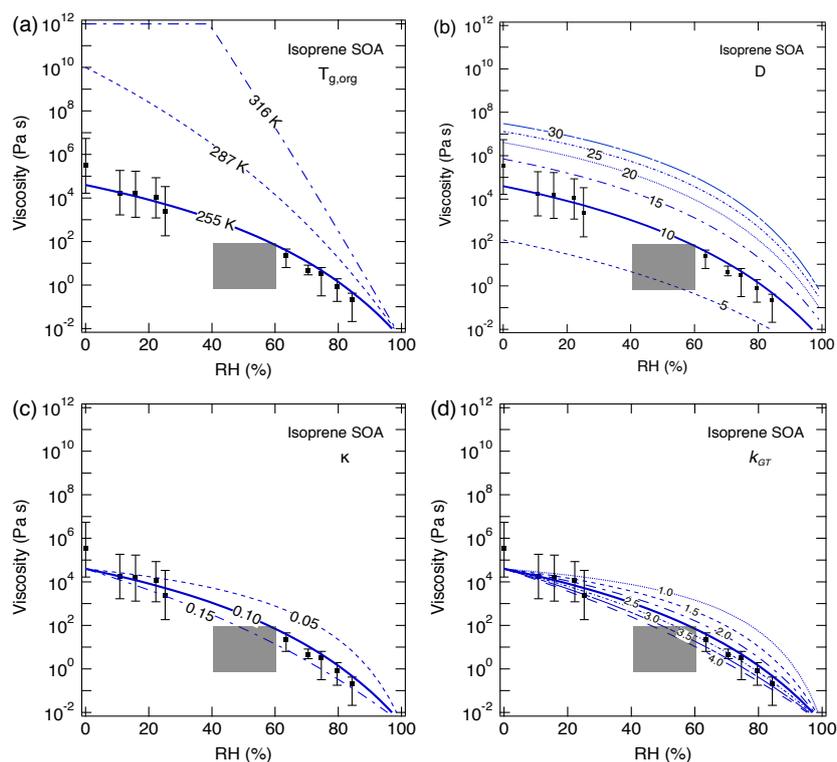
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Figure 4. Comparison of measured and predicted viscosity of (a) α -pinene SOA and (b) isoprene SOA at 295 K as a function of RH. The solid lines represent base simulations with the VTF equation, while the dotted line represents viscosity predicted using the WLF equation [parameters: glass transition temperature of dry SOA ($T_{g,org}$), fragility (D), hygroscopicity (κ) and Gordon-Taylor constant (k_{GT}): (a) 278.5 K, 0.1, 10 and 2.5; (b) 255 K, 0.1, 10 and 2.5. The shaded regions were determined by varying these parameters (a) upper (lower) limit: $T_{g,org} = 300$ K (278.5 K), $\kappa = 0.1$ (0.1), $D = 20$ (10), $k_{GT} = 2.5$ (2.0); (b) upper (lower limit): $T_{g,org} = 255$ K (255 K), $\kappa = 0.10$ (0.15), $D = 15$ (8), $k_{GT} = 2.5$ (4.0). Grayson et al. (2016) data in the panel (a) represents two different mass loadings ($121 \mu\text{g m}^{-3}$; $520 \mu\text{g m}^{-3}$). The gray box in panel (b) represents estimated viscosity based on bounce measurements of Bateman et al. (2015).



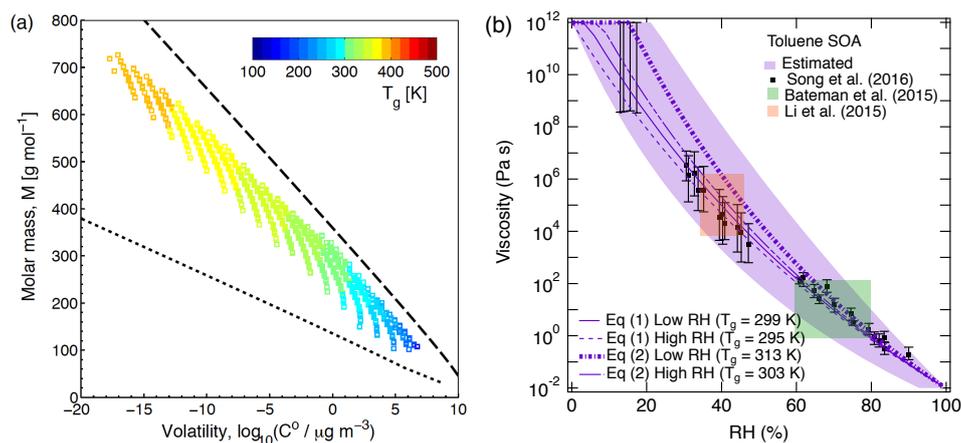
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832 **Figure 5.** Sensitivity calculations for viscosity of α -pinene SOA at 295 K as a function of RH
833 by varying: (a) glass transition temperature of dry SOA ($T_{g,org}$), (b) fragility (D), (c)
834 hygroscopicity (κ), and (d) Gordon-Taylor constant (k_{GT}).



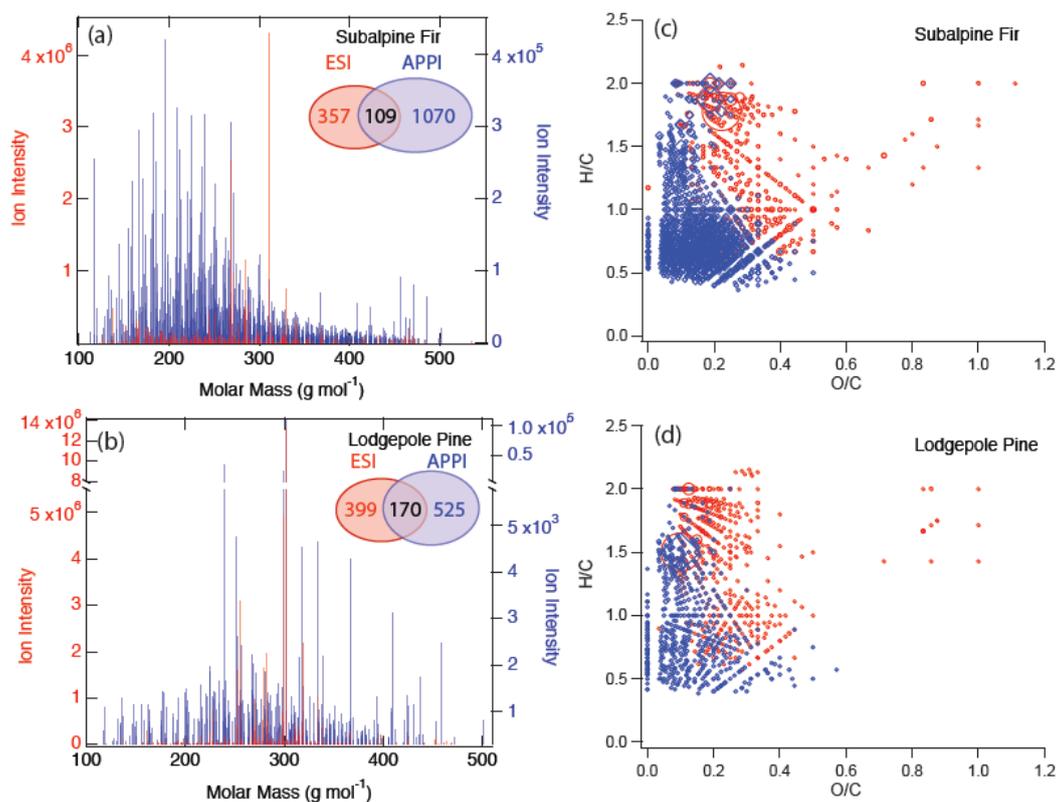
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Figure 6. Sensitivity calculations for viscosity of isoprene SOA at 295 K as a function of RH by varying: (a) glass transition temperature of dry SOA ($T_{g,org}$), (b) fragility (D), (c) hygroscopicity (κ), and (d) Gordon-Taylor constant (k_{GT}). Data points are measured viscosity by Song et al. (2015) and the gray box represents estimated viscosity based on bounce measurements of Bateman et al. (2015).

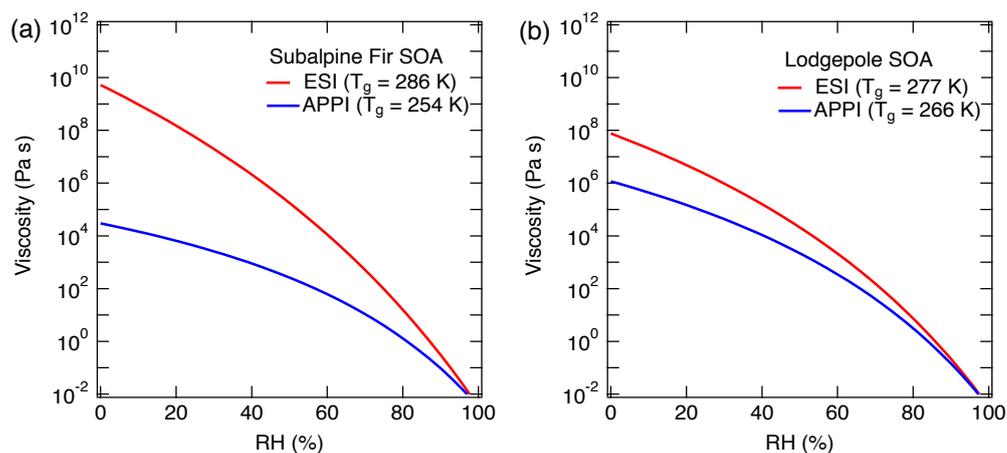


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844 **Figure 7.** (a) Molecular corridor of molar mass plotted against volatility of toluene SOA formed
 845 under dry conditions (Hinks et al., 2017) color-coded by glass transition temperature (T_g)
 846 estimated using Eq. (2). The upper dashed line indicates the low O:C bound of the molecular
 847 corridor (linear alkanes C_nH_{2n+2} with O:C = 0), and the lower dotted line indicates the high O:C
 848 bound (sugar alcohols $C_nH_{2n+2}O_n$ with O:C = 1). (b) Comparison of measured (markers) and
 849 modeled (lines) viscosity of toluene SOA at 295 K as a function of RH. Viscosities were
 850 calculated using fragility (D) of 13, the hygroscopicity (κ) of 0.25 and the Gordon-Taylor
 851 constant (k_{GT}) of 3.0 with different glass transition temperatures of dry SOA ($T_{g,org}$) as estimated
 852 using Eq. (1) or (2) under low and high RH conditions. The shaded regions were calculated by
 853 varying those parameters: $T_{g,org} = 313$ K (295 K), $\kappa = 0.20$ (0.25), $D = 13$ (10), $k_{GT} = 2.5$ (3.5) for
 854 the upper (lower) limit.
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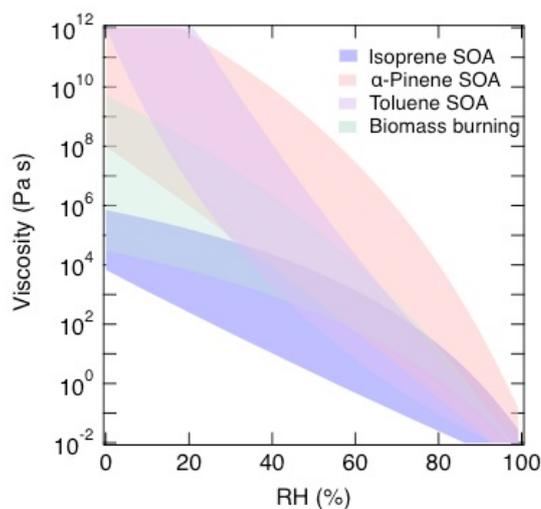


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857 **Figure 8.** Mass spectra of biomass burning organic particles collected from test burns of (a)
858 subalpine fir and (b) lodgepole pine as measured by high resolution mass spectrometry with
859 two ionization techniques: electron spray ionization (ESI, red) and atmospheric pressure
860 photoionization (APPI; blue). Numbers of elemental formulas identified by ESI (red), APPI
861 (blue) and both modes (black) are also specified. Van Krevelen plots of the compounds
862 identified by ESI (red) and APPI (blue) mode in BBOA from burning of (c) subalpine fir and (d)
863 lodgepole pine.
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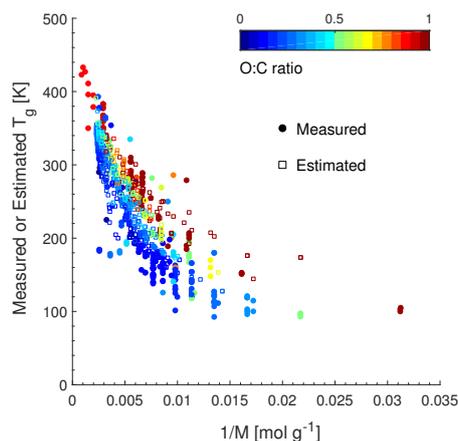
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Figure 9. Predicted viscosity for biomass burning particles of (a) subalpine fir and (b) lodgepole pine trees as measured by high resolution mass spectrometry with two ionization techniques: electrospray ionization (ESI, red) and atmospheric pressure photoionization (APPI; blue). $T_{g,org}$ are specified in the figure legend and other used parameters are fixed to $\kappa = 0.1$, $D = 10$, $k_{GT} = 2.5$.



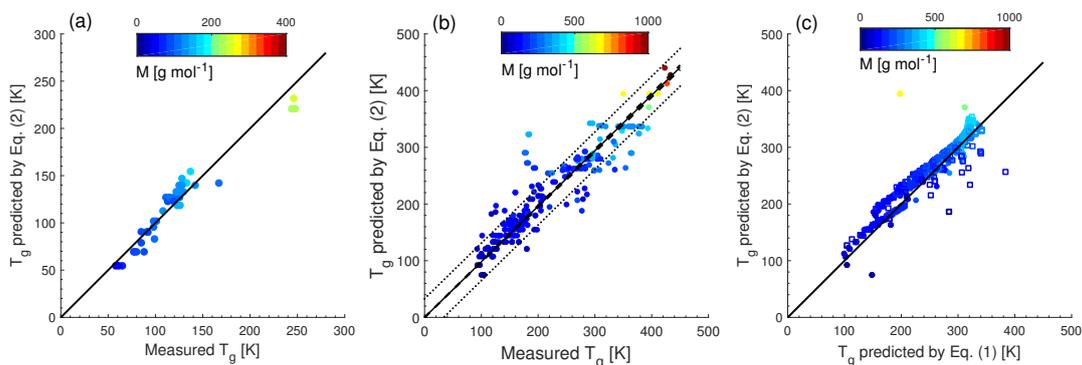
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Figure 10. Summary of predicted range of viscosity of α -pinene SOA (red), isoprene SOA (blue), toluene SOA (purple), and biomass burning particles (green).



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880 **Figure A1.** T_g of organic compounds as measured (circles) and estimated with the Boyer-
881 Kauszmann rule (squares) plotted against the inverse molar mass. The markers are color-coded by
882 atomic O:C ratio.
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886 **Figure A2.** Predicted T_g using Eq. (2) compared with (a) measured T_g for (a) CH and (b) CHO
887 compounds and (c) predicted T_g using Eq. (1) for CHO compounds. The solid line shows 1:1 line
888 and the dashed and dotted lines in the panel (c) show 68% confidence and prediction bands,
889 respectively. In panel (c), solid circle markers represent measured T_g and open square markers
890 represent estimated T_g by the Boyer-Kauszmann rule.