1	Predicting the glass transition temperature and viscosity of secondary				
2	organic material using molecular composition				
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17 Abstract:

18 Secondary organic aerosols (SOA) account for a large fraction of submicron particles in the 19 atmosphere. SOA can occur in amorphous solid or semi-solid phase states depending on chemical composition, relative humidity (RH), and temperature. The phase transition between 20 21 amorphous solid and semi-solid states occurs at the glass transition temperature (T_g) . We have recently developed a method to estimate $T_{\rm g}$ of pure compounds containing carbon, hydrogen, and 22 oxygen atoms (CHO compounds) with molar mass less than 450 g mol⁻¹ based on their molar 23 24 mass and atomic O:C ratio. In this study, we refine and extend this method for CH and CHO compounds with molar mass up to $\sim 1100 \text{ g mol}^{-1}$ using the number of carbon, hydrogen, and 25 26 oxygen atoms. We predict viscosity from the T_{g} -scaled Arrhenius plot of fragility (viscosity vs. 27 $T_{\rm g}/T$) as a function of the fragility parameter D. We compiled D values of organic compounds 28 from literature, and found that D approaches a lower limit of ~ 10 (+/- 1.7) as the molar mass 29 increases. We estimated viscosity of α -pinene and isoprene SOA as a function of RH by 30 accounting for hygroscopic growth of SOA and applying the Gordon-Taylor mixing rule, 31 reproducing previously published experimental measurements very well. Sensitivity studies were conducted to evaluate impacts of T_g , D, hygroscopicity parameter (κ), and the Gordon-Taylor 32 33 constant on viscosity predictions. Viscosity of toluene SOA was predicted using the elemental 34 composition obtained by high-resolution mass spectrometry (HRMS), resulting in a good 35 agreement with the measured viscosity. We also estimated viscosity of biomass burning particles 36 using the chemical composition measured by HRMS with two different ionization techniques: 37 electrospray ionization (ESI) and atmospheric pressure photoionization (APPI). Due to 38 differences in detected organic compounds and signal intensity, predicted viscosities at low RH 39 based on ESI and APPI measurements differ by 2-5 orders of magnitude. Complementary

40 measurements of viscosity and chemical composition are desired to further constrain RH41 dependent viscosity in future studies.

42

43 **1. Introduction**

44 Secondary organic aerosols (SOA) account for a large fraction of submicron particles in the 45 atmosphere and they play an important role in climate, air quality and public health (Goldstein 46 and Galbally, 2007; Jimenez et al., 2009). Traditionally, SOA particles were assumed to be liquid with dynamic viscosity η below 10² Pa s, but a number of recent studies have shown that 47 they can also adopt amorphous semi-solid ($10^2 \le \eta \le 10^{12}$ Pa s), or glassy solid ($\eta > 10^{12}$ Pa s) 48 49 states, depending on chemical composition and temperature (Zobrist et al., 2008; Koop et al., 50 2011; Huang et al., 2018; Reid et al., 2018). The phase state is also strongly affected by relative 51 humidity, as water can act as a plasticizer to lower viscosity (Mikhailov et al., 2009). Ambient 52 and laboratory-generated SOA particles have been observed to bounce off the smooth hard 53 surface of an inertial impactor at low RH, implying a non-liquid state (Virtanen et al., 2010; 54 Saukko et al., 2012; Bateman et al., 2015; Jain and Petrucci, 2015), whereas predominantly 55 biogenic SOA particles in the Amazon basin did not bounce off the impactor surface at high RH, 56 implying they are primarily liquid (Bateman et al., 2016). Upon dilution or heating, SOA 57 particles were observed to evaporate unexpectedly slowly (Cappa and Wilson, 2011; Vaden et 58 al., 2011), and recent modeling studies have evaluated the contributions of low diffusivity and 59 volatility to slow evaporation rates (Roldin et al., 2014; Yli-Juuti et al., 2017). Measurements of 60 viscosity of SOA bulk material derived from oxidation of α -pinene (Renbaum-Wolff et al., 2013; 61 Zhang et al., 2015; Hosny et al., 2016), limonene (Hinks et al., 2016), isoprene (Song et al.,

62 2015), and toluene (Song et al., 2016a) have confirmed that SOA particles adopt a wide range of63 viscosities.

The particle phase state has been shown to affect gas uptake and chemical transformation of 64 organic compounds due to kinetic limitations of bulk diffusion (Shiraiwa et al., 2011; Abbatt et 65 al., 2012; Kuwata and Martin, 2012; Zhou et al., 2013; Slade and Knopf, 2014; Arangio et al., 66 67 2015; Davies and Wilson, 2015; Wang et al., 2015; Berkemeier et al., 2016; Marshall et al., 68 2016; Liu et al., 2018; Pratap et al., 2018; Zhang et al., 2018). Molecular motion can be hindered 69 in a highly viscous matrix, slowing down photochemical reactions in particles (Lignell et al., 70 2014; Hinks et al., 2016). Water diffusion can be still fast even in an amorphous solid matrix 71 under room temperature, but it can be hindered significantly under low temperatures (Mikhailov 72 et al., 2009; Zobrist et al., 2011; Bones et al., 2012; Berkemeier et al., 2014; Price et al., 2014), 73 affecting homogeneous vs. heterogeneous ice nucleation pathways (Murray et al., 2010; Wang et 74 al., 2012a; Wang et al., 2012b; Baustian et al., 2013; Schill and Tolbert, 2013; Berkemeier et al., 75 2014; Schill et al., 2014; Lienhard et al., 2015; Ignatius et al., 2016; Knopf et al., 2018). Despite 76 the substantial implications of the SOA particle phase state, its effects on gas-particle 77 interactions have not yet been considered explicitly in current climate and air quality models 78 (Shrivastava et al., 2017).

Partitioning of semi-volatile compounds into viscous particles may result in kineticallylimited growth in contrast to quasi-equilibrium growth (Perraud et al., 2012; Shiraiwa and Seinfeld, 2012; Booth et al., 2014; Zaveri et al., 2014; Mai et al., 2015; Liu et al., 2016), which also affects the evolution of particle size distribution upon SOA growth (Shiraiwa et al., 2013; Zaveri et al., 2018). Chamber experiments probing mixing timescales of SOA particles derived by oxidation of various precursors such as isoprene, terpene, and toluene have observed strong

kinetic limitations at low RH, but not at moderate and high RH (Loza et al., 2013; Ye et al., 2016; Ye et al., 2018). Gorkowski et al. (2017) did not observe significant diffusion limitations for glycerol and squalene in α -pinene SOA. Quasi-equilibrium versus kinetically-limited or nonequilibrium SOA growth remains an open issue and warrants further investigations.

89 Group contribution methods have been used to predict the viscosities of pure compounds 90 when the functionality and molecular structure are known (Sastri and Rao, 1992; Rothfuss and 91 Petters, 2017). Song et al. (2016) showed that estimations from group contribution approaches 92 combined with either nonideal or ideal mixing reproduced the RH-dependent trends particularly well for the alcohol, di-, and tricarboxylic acid systems with viscosity of up to 10^4 Pa s. By 93 94 contrast, model calculations overestimated the viscosity of more viscous compounds including 95 mono-, di-, and trisaccharides by many orders of magnitude (Song et al., 2016b). A recent study 96 compiled viscosity of organic compounds with atmospherically relevant functional groups, 97 investigating the influence of the number and location of functional groups on viscosity 98 (Rothfuss and Petters, 2017). These studies provide important insights in estimating the viscosity 99 of individual organic compounds.

Particle phase state can be characterized by a glass transition temperature (T_g) , which is a 100 characteristic temperature representing a non-equilibrium phase transition from a glassy solid 101 102 state to a semi-solid state as the temperature increases (Koop et al., 2011). Recently, we have developed a method to estimate T_{g} of pure organic compounds comprised of carbon, hydrogen, 103 and oxygen (CHO compounds) with molar mass less than 450 g mol⁻¹ based on their molar mass 104 105 and atomic O:C ratio (Shiraiwa et al., 2017). It has been applied successfully in a global chemistry climate model to predict T_g and the phase state of atmospheric SOA, which indicated 106 107 that SOA particles are mostly liquid or semi-solid in the planetary boundary layer, while they

should be glassy in the middle and upper troposphere (Shiraiwa et al., 2017). A recent study
provided a consistent result, suggesting that mixing timescales of organic molecules within SOA
are often < 1 h in a global planetary boundary layer (Maclean et al., 2017).

111 It has been shown that SOA particles contain oligometric compounds with molar masses higher than 450 g mol⁻¹ (Gao et al., 2004; Tolocka et al., 2004; Nizkorodov et al., 2011; Nozière 112 113 et al., 2015), which makes the previously developed parameterization incomplete. In this study, 114 we extend the parameterization of T_g to higher molar mass compounds, and apply it to high-115 resolution mass spectrometry data for toluene SOA and biomass burning particles. The 116 Arrhenius approach and the Gordon-Taylor mixing rules were applied to estimate viscosity of 117 SOA bulk materials to compare with the literature reported viscosity measurements. This method 118 will be useful for estimations of viscosity of organic particles, for which high-resolution mass 119 spectra are available. It can also be applied in global or regional models to evaluate impacts of 120 the particle phase state on the role of SOA in climate and air quality.

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122 **2. Parameterization development**

123 **2.1 Glass transition temperature**

Figure 1a shows the dependence of T_g on the molar mass (*M*) of organic compounds. Solid markers represent measured T_g of 258 CHO compounds (Koop et al., 2011; Dette et al., 2014; Rothfuss and Petters, 2017), while open markers represent 654 CHO compounds in SOA (Shiraiwa et al., 2014). Markers are color-coded by atomic O:C ratio. Their melting points (T_m) were estimated by the Estimation Programs Interface (EPI) Suite software version 4.1 (US-EPA, 2012) and their T_g were estimated using the Boyer-Kauzmann rule: $T_g = g \cdot T_m$ with g = 0.7 (Koop et al., 2011; Shiraiwa et al., 2017). This rule can provide good estimates of T_g , as has been

131 validated in previous work (Koop et al., 2011) and also shown in Fig. A2(a). A subset of data 132 shown in Figure 1 was originally published in Shiraiwa et al. (2017) for compounds with M < 1450 g mol⁻¹. This version of the figure has been updated to include a number of experimentally 133 measured T_g values of larger compounds with M up to 1153 g mol⁻¹, including aliphatic 134 135 compounds containing OH and/or COOH groups. Specifically, data for 76 aliphatic alcohols, 39 136 carbohydrates and their derivatives, 4 carboxylic acids, and 4 hydroxy acids, as compiled by 137 Rothfuss and Petters (2017), have been added to Figure 1. Eight of these compounds are carbohydrates with M > 450 g mol⁻¹. These updates are critical for reliable parameterization of $T_{\rm g}$ 138 based on M. When M increases above ~500 g mol⁻¹, the slope of T_g decreases, making it 139 140 challenging to extrapolate the low-M data from the original Shiraiwa et al. (2017) study to higher M values. When M increases to ~1000 g mol⁻¹, the corresponding $T_{\rm g}$ appears to level at around 141 142 420 K.

143 Such dependence on M has been described for polymers with the Fox-Flory equation: $T_{\rm g}(M) = T_{\rm g,\infty} - \frac{K_{\rm m}}{M}$ (Fox Jr and Flory, 1950), where $K_{\rm m}$ is a constant and $T_{\rm g,\infty}$ is the asymptotic 144 value of $T_{\rm g}$ specific to the polymer. We conducted a literature search and found that most of the 145 reported $T_{g,\infty}$ values fell below ~500 K (Fox Jr and Flory, 1950; Onder et al., 1972; Montserrat 146 147 and Colomer, 1984; Polymer handbook, 1999; Papadopoulos et al., 2004; Matsushima et al., 148 2017). The Fox-Flory equation works very well for high molar mass compounds and is also 149 generally applicable to smaller compounds (Koop et al., 2011), as supported by an approximately 150 linear dependence of T_g on the inverse molar mass in Fig. A1(a). Figure 1b plots the values of T_g 151 as a function of the atomic O:C ratio of organic molecules. Figures 1a and 1b clearly 152 demonstrate that $T_{\rm g}$ depends primarily on the molar mass with a weak dependence on the atomic 153 O:C ratio.

A parameterization for T_g calculation based on the molar mass and atomic O:C ratio was developed in our recent work, which is applicable to CH and CHO compounds with M < 450 g mol⁻¹ (Shiraiwa et al., 2017):

157
$$T_{\rm g} = A + BM + C^2 M^2 + D (O:C) + E M (O:C)$$
 (1)

where A = $-21.57 (\pm 13.47) [K]$, B = $1.51 (\pm 0.14) [K \mod g^{-1}]$, C = $-1.7 \times 10^{-3} (\pm 3.0 \times 10^{-4}) [K \mod^2 g^{-2}]$, D = $131.4 (\pm 16.01) [K]$ and E = $-0.25 (\pm 0.085) [K \mod g^{-1}]$, respectively. These values were obtained by fitting the measured T_g of 179 CH and CHO compounds with M < 450 g mol⁻¹ with multi-linear least squares analysis. Note that application of Eq. (1) may provide unreasonable T_g values for compounds with M > 500 g mol⁻¹ because it does not account for the strong curvature in the T_g vs. M dependence shown in Figure 1a.

In this study we have developed an improved parameterization to predict T_g of CH and CHO compounds using the number of carbon (n_c), hydrogen (n_H), and oxygen (n_O) that can also be applied to higher molar mass compounds. Motivated by a good correlation between T_g and volatility (Fig. 1a in Shiraiwa et al., (2017)), we use an equation with a similar formulation to the equation used to predict the saturation mass concentration or volatility (Donahue et al., 2011; Li et al., 2016):

170
$$T_{\rm g} = (n_{\rm C}^0 + \ln(n_{\rm C})) b_{\rm C} + \ln(n_{\rm H}) b_{\rm H} + \ln(n_{\rm C}) \ln(n_{\rm H}) b_{\rm CH} + \ln(n_{\rm O}) b_{\rm O} + \ln(n_{\rm C}) \ln(n_{\rm O}) b_{\rm CO}$$
(2)

171 where $n_{\rm C}^0$ is the reference carbon number, $b_{\rm C}$, $b_{\rm H}$ and $b_{\rm O}$ denote the contribution of each atom to 172 $T_{\rm g}$, and $b_{\rm CH}$ and $b_{\rm CO}$ are coefficients that reflect contributions from carbon-hydrogen and carbon-173 oxygen bonds, respectively. These values were obtained by fitting the measured $T_{\rm g}$ of 42 CH 174 compounds and 258 CHO compounds with multi-linear least squares analysis with 68% 175 prediction and confidence intervals. The best-fit parameters are summarized in Table 1.

176 Note that the evaluation dataset used to derive Eq. (2) contains CH compounds with $M < \infty$ 177 260 g mol⁻¹ (see Fig. A2b for comparison of measured and predicted T_g). Thus, the application of Eq. (2) to higher molar mass compounds may require further investigations when measured $T_{\rm g}$ 178 179 for higher molar mass compounds becomes available. We plan to continue to refine our method 180 as additional glass transition data on high molar mass compounds become available. Figure 1c shows that the T_g values predicted using Eq. (2) are in good agreement with the T_g values 181 182 measured in experiments (see also Fig. A1(b)) or estimated by the Boyer-Kauzmann rule as indicated by the high correlation coefficient of 0.95. $T_{\rm g}$ of individual compounds can be 183 184 predicted within ± 21 K as indicated by the prediction band (dotted lines in Fig. 1c); however, 185 this uncertainty may be much smaller for multicomponent SOA mixtures under ideal mixing 186 conditions as indicated in the confidence band (dashed lines, almost overlapping with the 1:1 187 line).

188 These results are noteworthy given that the parameterization (Eq. 2) does not consider either explicit molecular structures or functional groups. Previous studies have shown that $T_{\rm g}$ can 189 190 be especially sensitive to the number of OH groups, which interact strongly through hydrogen bonding. For example, Nakanishi et al., (2011) found a direct relationship between T_g and the 191 192 number of hydroxyl groups in a molecule for sugar alcohols; $T_{\rm g}$ increases as the number of OH 193 groups increases. They reported that the correlation between $T_{\rm g}$ and the number of OH groups 194 was much stronger than the correlation between $T_{\rm g}$ and the number of carbons in a molecule. 195 Such a trend is implicitly included in Eq. (1) and (2), which contain the O:C ratio and number of 196 oxygen atoms as parameters, respectively. Recently, Rothfuss and Petters (2017) showed an approximately linear relationship between the number of OH groups and T_{g} for compounds with 197 198 up to eight OH groups. Grayson et al. (2017) showed that addition of hydroxyl functional groups

199 increases viscosity, a conclusion supported by both the experimental data and quantitative 200 structure-property relationship model. The correlation between $T_{\rm g}$ and the number of carbon 201 atoms is consistent with the free volume theory, in which molecular motion is restricted by the 202 difference between the space required for a molecule to vibrate versus the space in which the 203 molecule resides (i.e., the free volume) (White and Lipson, 2016). The correlation between $T_{\rm g}$ 204 and the number of OH groups is more consistent with the topological constraint theory, where 205 the primary influence is the three dimensional structure of the molecule as determined by 206 molecular bonds and hydrogen-bonding networks (Nakanishi and Nozaki, 2011; van der Sman, 207 2013). Future experiments targeting more comprehensive T_g data, especially for higher molar mass compounds, would lead to further refinements of our T_{g} parameterizations. 208

209 Comparing Eq. (1) and (2), the two parameterizations give similar performance for compounds with M < 450 g mol⁻¹ as shown in Fig. A2c. The statistical measures of correlation 210 211 coefficient (R), mean bias (MB), and root mean square error (RMSE) are 0.93, -6.45 K, and 212 25.64 K, respectively, for the performance of Eq. (1), while for Eq. (2), they are 0.95, 3.15 K, 213 and 21.11 K, respectively. It should be noted again that Eq. (1) cannot be used to predict T_g for compounds with M > 450 g mol⁻¹. For example, T_g of stachyose (M = 667 g mol⁻¹) predicted by 214 215 Eq. (1) is 198 K, while that by Eq. (2) is 394 K, which agrees much better with the measured mean T_g of 396 K (Rothfuss and Petters, 2017). Eq. (2) is more flexible than Eq. (1) and can be 216 217 potentially expanded to include compounds containing hetero-atoms (e.g., nitrogen or sulfur), 218 once substantial sets of experimental values of T_g for such compounds become available. 219 Regarding the application in air quality and climate models, Eq. (1) can be applied in the 220 volatility basis set (VBS) (Donahue et al., 2006; Donahue et al., 2011) and the molecular corridor approach (Shiraiwa et al., 2014; Li et al., 2016) to predict the $T_{\rm g}$ of SOA particles 221

222 (Shiraiwa et al., 2017), while the new parameterization may be suitable for coupling with the 223 statistical oxidation model which characterizes the SOA evolution as a function of $n_{\rm C}$ and $n_{\rm O}$ 224 (Cappa and Wilson, 2012; Jathar et al., 2015).

225 These parameterizations (Eqs. 1, 2) calculate T_g based on the elemental composition of 226 organic compounds. SOA particles contain a number of organic compounds as well as a variable 227 amount of liquid water, which has low T_g (136 K) and can act as a plasticizer (Mikhailov et al., 228 2009; Koop et al., 2011). Under humid conditions, SOA particles take up water by hygroscopic growth in response to RH, lowering T_g and viscosity of SOA particles. Estimations of T_g for 229 230 SOA-water mixtures were discussed by Shiraiwa et al. (2017), who applied the Gordon-Taylor 231 equation validated for a wide range of mixtures of organics, polymer, and water (Roos, 1993; 232 Hancock and Zografi, 1994; Zobrist et al., 2008; Dette et al., 2014; Dette and Koop, 2015). 233 Briefly, T_g of mixtures of SOA compounds under dry conditions ($T_{g,org}$) were calculated assuming the Gordon-Taylor constant (k_{GT}) of 1 (Dette et al., 2014): $T_{g,\text{org}} = \sum_i w_i T_{g,i}$, where w_i is 234 235 the mass fraction of organic compound *i*, which can be derived using mass concentrations of SOA products. The Gordon-Taylor equation can also be applied to calculate T_g of organic-water 236 237 mixtures considering the mass fraction of organics (w_{org}) in SOA particles (Koop et al., 2011):

238
$$T_{\rm g}(w_{\rm org}) = \frac{(1 - w_{\rm org})T_{\rm g,w} + \frac{1}{k_{\rm GT}}w_{\rm org}T_{\rm g,org}}{(1 - w_{\rm org}) + \frac{1}{k_{\rm GT}}w_{\rm org}}$$
(3)

239 w_{org} can be calculated using the mass concentrations of water $(m_{\text{H}_2\text{O}})$ and SOA (m_{SOA}) as $w_{\text{org}} =$ 240 $m_{\text{SOA}} / (m_{\text{SOA}} + m_{\text{H}_2\text{O}})$. $m_{\text{H}_2\text{O}}$ can be estimated using the effective hygroscopicity parameter (κ) 241 (Petters and Kreidenweis, 2007):

242
$$m_{\rm H2O} = \frac{\kappa \rho_{\rm w} m_{\rm SOA}}{\rho_{\rm SOA} \left(\frac{1}{a_W} - 1\right)} \tag{4}$$

The density of water (ρ_w) is 1 g cm⁻³, the density of SOA particles (ρ_{SOA}) is assumed to be 1.2 g cm⁻³ (Kuwata et al., 2012), m_{SOA} is the total mass concentrations of SOA, and a_w is the water activity calculated as $a_w = RH/100$. Pajunoja et al. (2015) found that water uptake in subsaturated conditions is inhibited until RH is high enough for dissolution of water in SOA particles with relatively low O:C ratios. As oxidation of SOA increases, solubility of water increases and dissolution occurs at lower RH values. In both cases, the use of subsaturated hygroscopicity measurements was supported.

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262

251 **2.2 Viscosity**

252 Temperature dependence of viscosity (η) was predicted using the modified Vogel-253 Tammann-Fulcher (VTF) equation (Angell, 1991):

$$\eta = \eta_{\infty} e^{\frac{T_0 D}{T - T_0}} \tag{5}$$

where η_{∞} is viscosity at infinite temperature; T_0 is the Vogel temperature; T is the ambient temperature. The fragility parameter, D, characterizes how rapidly the dynamics of a material slow down as T approaches T_g , reflecting to what degree the temperature dependence of the viscosity deviates from Arrhenius behavior. When T is close to T_g ($T_g/T \approx 1$), smaller D values indicate that viscosity is sensitive to temperature change (fragile behavior); while larger D values indicate that viscosity is less sensitive to temperature change (strong or Arrhenius behavior).

261 Assuming
$$\eta_{\infty} = 10^{-5}$$
 Pa s (Angell, 1991):

$$\log \eta = -5 + 0.434 \frac{T_0 D}{T - T_0} \tag{6}$$

263 When $T = T_g$, $\eta = 10^{12}$ Pa s, which leads to (Angell, 1991; Angell, 2002):

264
$$T_0 = \frac{39.17 \, T_g}{D+39.17} \tag{7}$$

As can be seen in Eq. (5), both T_g and D are required to calculate η from Eq. (4) at a given temperature.

Figure 2 shows the T_g -scaled Arrhenius plot of fragility (viscosity versus T_g/T) referred to 267 as an Angell plot (Angell, 1995). D values of organic compounds are typically in the range of 268 269 \sim 5–30 (Angell, 1997). To estimate D values that could be applied to SOA compounds, we 270 compiled measured fragility values. Fragility was often measured in the form of the fragility steepness index (m), which represents the slope of the Arrhenius plot at the point where $T = T_g$ 271 272 (Boehmer et al., 1993). Compounds with lower m exhibit higher D values, indicating stronger 273 glass formers. The measured m of 95 organic compounds are included in the Supplement. m can 274 be converted to D using the following equation (see the full derivation of this equation in 275 Appendix A):

276

$$D = \frac{665.89}{m - 17} \tag{8}$$

Figure 3 shows the measured *D* as a function of (a) molar mass and (b) atomic O:C ratio of organic molecules. The molar mass exerts a stronger effect on fragility, while there is little dependence of *D* on the O:C ratio. As molar mass increases, *D* approaches a lower limit of 10.3 (± 1.7) , consistent with the value of 10 used in our recent study (Shiraiwa et al., 2017). To evaluate the impact of the variations of *D* on viscosity prediction, sensitivity calculations were conducted as described in Sect. 3.

Besides the VTF equation, another commonly used equation for describing the temperature dependence of viscosity is the Williams-Landel-Ferry (WLF) equation: $\log \frac{\eta(T)}{\eta(T_g)} = \frac{-C_1(T-T_g)}{C_2+(T-T_g)}$, where empirical parameters C_1 and C_2 are adopted as 17.44 and 51.6 K, respectively (Williams et al., 1955; Schill and Tolbert, 2013; Wang et al., 2015). The two equations are

287 mathematically equivalent, both defined with respect to a reference temperature, and their parameters are related through $C_1 = \frac{DT_0}{2.303(T_g - T_0)}$ and $C_2 = T_g - T_0$. For the WLF equation, T_g is 288 289 the reference temperature and there is a linear dependence assumed between temperature and 290 free volume (O'Connell and McKenna, 1999; Huang and McKenna, 2001; Metatla and Soldera, 291 2007). For the VTF equation, the reference is the Vogel temperature (T_0) -a hypothetical 292 temperature at which all non-vibrational motion ceases and viscosity becomes infinite and the 293 theoretical foundation of the VTF equation includes both thermodynamic and kinetic 294 considerations (O'Connell and McKenna, 1999; Huang and McKenna, 2001; Metatla and 295 Soldera, 2007). The calculations of viscosity in this study are based mainly on the VTF equation 296 and the difference between calculated results from the two equations will be briefly discussed in 297 the following section.

298

3. Comparison of predicted viscosity with measurements

300 **3.1. SOA formed from** α **-pinene and isoprene**

301 The purpose of this section is to demonstrate that viscosity of SOA material can be predicted over a broad range of RH values from four parameters: T_g of dry SOA ($T_{g,org}$), fragility 302 303 (D), hygroscopicity (κ), and the Gordon-Taylor constant for mixing SOA and water (k_{GT}). 304 Viscosity of α -pinene SOA has been measured as a function of RH by several groups using 305 multiple experimental techniques as shown in Fig. 4(a) (Abramson et al., 2013; Renbaum-Wolff 306 et al., 2013; Kidd et al., 2014; Pajunoja et al., 2014; Bateman et al., 2015; Zhang et al., 2015; 307 Grayson et al., 2016). The wide range of experimentally measured viscosities reported for α -308 pinene SOA, particularly from 30-60% RH is most likely a consequence of the different 309 experimental approaches, mass loadings and O:C ratios for each experiment. For instance,

Grayson et al. (2016) used mass loadings of 121 to 14000 µg m³ and observed that viscosity 310 311 decreased as mass loading increased. Higher mass loadings would lead to greater partitioning of 312 semi-volatile and lower molar mass compounds into the particle phase, which would lead to the decrease of $T_{\rm g}$ and viscosity of the resulting SOA mixture. They concluded that their results 313 314 should be considered a lower limit for viscosity of α -pinene SOA in the atmosphere. It should 315 also be noted that the viscosity measurements from Renbaum-Wolff et al. (2013) were for the 316 water-soluble portion of the SOA. These datasets suggest that viscosity of α -pinene SOA approaches very high values ($\sim >10^8$ Pa s) below 20-30% RH and decreases with an increase in 317 318 RH reaching a value of ~10 Pa s at 80% RH. As can be seen in Fig. 4(b), isoprene SOA is less viscous with $\eta < 10^6$ Pa s even under dry conditions, undergoing a phase transition from a semi-319 320 solid phase to a liquid phase at ~55% RH (Bateman et al., 2015; Song et al., 2015).

321 The solid lines with the shaded areas in Figure 4 are viscosity values predicted using $T_{g,org}$, D, κ , k_{GT} . $T_{g,org}$ values were adopted by Berkemeier et al. (2014) who estimated $T_{g,org}$ with 322 323 the Boyer-Kauzmann rule using the melting point of representative SOA oxidation products. Note that Eq. (1) or (2) were not used to estimate $T_{g,org}$, which should be done in future studies 324 325 by obtaining their elemental composition using high resolution mass spectrometry. For α -pinene, 326 $T_{g,org}$ was assumed to be 278 K corresponding to an O:C ratio of 0.5 (Berkemeier et al., 2014), 327 which is a typical O:C ratio of α -pinene SOA (Aiken et al., 2008; Chen et al., 2011; Putman et 328 al., 2012).

The $T_{g,org}$ selected for isoprene SOA was 255 K, corresponding to the O:C ratio of 0.6. Although no measurements of the O:C ratio for the experimental isoprene SOA data were reported, Song et al. (2015) estimated O:C of 0.64-1.1 based on literature values. As O:C ratios are useful in estimating $T_{g,org}$, we encourage the measurement of the O:C ratio of SOA when

333 conducting viscosity measurements. In contrast to α -pinene SOA, there are limited viscosity 334 measurements for isoprene SOA. While the predicted viscosity is consistent with the 335 experimental data, comparison of our model predictions to additional measurements is strongly 336 recommended. Song et al. (2015) prepared their samples in a potential aerosol mass (PAM) 337 reactor while those investigated by Bateman et al. (2015) were generated in a smog chamber. It 338 has been suggested that under ambient conditions, the majority of isoprene-derived SOA can be 339 derived through heterogeneous interactions with acidic sulfate particles forming oligomers (Lin 340 et al., 2013; Surratt et al., 2010; Gaston et al., 2014), which may increase viscosity. Further 341 studies are warranted to compare laboratory-generated and ambient isoprene SOA, and to 342 investigate the effect of the acidic seed on the viscosity.

343 For both α -pinene and isoprene SOA, D was set to 10 based on the analysis presented in 344 Fig. 3(a). κ was set to 0.1 based on field and laboratory measurements (Gunthe et al., 2009; 345 Lambe et al., 2011b; Pajunoja et al., 2014; Petters et al., 2017) and k_{GT} was assumed to be 2.5 346 (Zobrist et al., 2008; Koop et al., 2011). Using these parameters, the predicted viscosities match 347 well the magnitude and the RH-dependence of the measured viscosity of α -pinene and isoprene 348 SOA. Figure 4 also shows predicted viscosities (dotted lines) using the WLF equation, which 349 shows similar values as the VTF equation, but slightly underestimates the viscosity of α -pinene 350 SOA at low RH and overestimates the viscosity of isoprene SOA at high RH.

Sensitivity studies were conducted to examine the effects of $T_{g,org}$, D, κ and k_{GT} , on the calculated viscosity. In these studies, $T_{g,org}$ of α -pinene and isoprene SOA were varied within 229-328 K and 255-316 K, respectively, representing $T_{g,org}$ of different oxidation states (Berkemeier et al., 2014). D was varied between 5 and 30, which is the range characteristic for organic compounds (see Fig. 3a). κ of 0.05-0.15 were used for α -pinene and isoprene SOA (Lambe et al., 2011b; Pajunoja et al., 2015). For the Gordon-Taylor constant, values of 2.5±1.5
were considered (Zobrist et al., 2008; Koop et al., 2011; Dette et al., 2014; Dette and Koop,
2015).

359 The effect of varying each parameter on the calculated viscosity of α -pinene SOA is 360 illustrated in Fig. 5. Variations of ± 50 K in $T_{g,org}$ result in 3-6 orders of magnitude difference in calculated values at dry conditions, indicating that $T_{g,org}$ is a critical parameter for viscosity 361 362 estimations. Decreasing D from 10 to 5 led to a decrease of calculated values by more than one 363 order of magnitude. The calculated results were within the upper limit of measurements when 364 increasing D from 10 to 20, and the predicted values were only slightly enhanced when further 365 increasing D from 20 to 30. Calculated values with variations in κ from 0.05 to 0.15 and $k_{\rm GT}$ 366 from 1.0 to 4.0 were all within the measured ranges.

367 For isoprene SOA, an increase of $T_{g,org}$ to 287 K, which represents a higher oxidation 368 state (Berkemeier et al., 2014), led to calculated values to be several orders of magnitude higher 369 than the upper limit of measurements (Fig. 6a). When $T_{g,org}$ reaches 316 K, isoprene SOA can 370 occur as a solid for RH lower than ~40%. Compared to α -pinene SOA, a variation in D has a 371 larger effect on the calculated viscosity (Fig. 6b). For a range of 5 - 30 for D, calculations with 372 the D value of 10 agreed well with the measurements, while other D values resulted in calculated viscosity outside of the measured ranges. Figures 6c and 6d show that decreasing κ and $k_{\rm GT}$ 373 374 below the reference values, the predictions overestimate the measured η by one or two orders of magnitude. The latter is most evident at RH > 60%, where the calculated values were higher than 375 376 the upper limit of measurements. Modeling results with κ and $k_{\rm GT}$ increasing to 0.15 and 4.0, 377 respectively, were within the lower limit of measurements.

378 The above comparison between the measured and predicted viscosity demonstrates that 379 the method described in this study can reproduce reasonably well the measured RH-dependent 380 viscosity of SOA formed from α -pinene and isoprene. The sensitivity calculations showed that $T_{g,org}$ contributed the most to the uncertainty in the viscosity estimates. Previous studies have 381 382 shown that the experimental conditions such as particle mass concentrations (Grayson et al., 383 2016) and RH upon SOA formation (Kidd et al., 2014; Hinks et al., 2018) can impact chemical 384 composition of SOA and hence the phase state and viscosity. Further efforts to constrain the 385 uncertainties are needed both in experiments and parameterizations.

386

387 **3.2. SOA formed from toluene**

388 In this and the following sections, we examine the feasibility of calculating the value of $T_{g,org}$ from mass spectrometry data on SOA. Hinks et al. (2017) measured the elemental 389 390 composition of toluene SOA using nanospray desorption electrospray ionization high-resolution 391 mass spectrometry (nano-DESI-HRMS) (Roach et al., 2010a, b). Toluene SOA were formed by OH photooxidation in an aerosol smog chamber at <2% RH (mass loading = 23 μ g m⁻³) and 75% 392 RH (mass loading = 8 μ g m⁻³) to investigate the effect of RH on the chemical composition of 393 394 toluene SOA formed under low-NO_x conditions. Measurements revealed a significant reduction 395 in the fraction of oligomers present in toluene SOA generated under high RH conditions 396 compared to SOA generated under low RH conditions (Hinks et al., 2017). The detected molar mass of individual oxidation products spanned a range of 102 - 570 g mol⁻¹ at high RH, which 397 increased up to 726 g mol⁻¹ at low RH. 398

Figure 7(a) shows the interdependence of glass transition temperature, volatility, and molar mass of the detected toluene SOA compounds. Glass transition temperatures were

401 calculated using Eq. (2). Saturation mass concentrations or volatilities of detected compounds 402 were estimated from the elemental composition by using the parameterization of Li et al. (2016). 403 The analysis is based on the molecular corridor approach—a two-dimensional framework of 404 volatility and molar mass of SOA components constrained by boundary lines of low and high 405 atomic O:C ratio, corresponding to *n*-alkanes (C_nH_{2n+2} , O:C = 0) and sugar alcohols ($C_nH_{2n+2}O_n$, 406 O:C = 1), respectively (Shiraiwa et al., 2014; Li et al., 2016). The toluene SOA constituents are 407 well constrained by the molecular corridor and $T_{\rm g}$ are higher for compounds with higher molar 408 mass and lower volatility.

Eq. (1) was used to calculate T_g for individual compounds with $M < 450 \text{ g mol}^{-1}$, while 409 excluding compounds with molar mass higher than 450 g mol⁻¹. This approach was deemed 410 411 reasonable as such high molar mass compounds account for < 10% of all toluene SOA products 412 formed at low RH, and for < 2% formed at high RH. Eq. (2) was used to calculate T_g for all the detected compounds. T_g of dry toluene SOA ($T_{g,org}$) was then computed using the Gordon-Taylor 413 414 approach with $k_{\text{GT}} = 1$ (Sect. 2.1). The relative mass concentrations of individual components 415 were assumed to be proportional to their relative abundance in the nano-DESI-HRMS spectrum. 416 This assumption has a number of caveats (Bateman et al., 2012; Nguyen et al., 2013), and as we 417 will see below, it results in deviations between the predicted and measured viscosity. Table 2 418 summarizes the results of such calculations, showing that the $T_{g,org}$ by Eq. (1) – excluding high molar mass compounds – is about 10 K lower as compared to $T_{g,org}$ by Eq. (2). $T_{g,org}$ at low RH is 419 420 predicted to be higher than $T_{g,org}$ at high RH, which results from a lower abundance of high molar 421 mass compounds observed at high RH. This trend is consistent with Kidd et al. (2014), who 422 showed that SOA material formed under dry conditions is more viscous than that formed under 423 wet conditions.

424 Figure 7(b) shows the predicted viscosity of toluene SOA as a function of RH, as 425 compared to the measured viscosity of toluene SOA formed in an oxidation flow reactor at 13% 426 RH (Song et al., 2016a). Indirect viscosity measurements are also included in shaded boxes 427 (Bateman et al., 2015; Li et al., 2015). Lines with shaded areas are calculated viscosities using 428 $T_{g,org}$ as described above. κ was assumed to be 0.25 based on laboratory measurements (Lambe et 429 al., 2011a; Hildebrandt Ruiz et al., 2015). To achieve good fit, D was set to 13 and k_{GT} was 430 assumed to be 3.0 (Dette et al., 2014). Estimations with Eq. (1) match the measured viscosity 431 values very well over the entire RH range. Predictions with Eq. (2) overestimated the 432 measurements by one or two orders of magnitude at moderate RH between 30% and 50%, while 433 they agreed with the measurements derived at $RH \ge 60\%$ and at the dry conditions.

434 There are several possible reasons for the difference between the measurements and 435 predictions. First, the relative abundance of high molar mass compounds observed in HRMS 436 measurements may be overestimated, as high molar mass compounds tend to have higher (yet 437 generally unknown) ionization efficiencies compared to lower molar mass compounds. Second, 438 the nano-DESI-HRMS analysis of toluene SOA was limited to m/z range of 100 -1000 (Hinks et 439 al., 2017). It is possible that some SOA products with lower molar mass were present in particles but not detected, which would lead to an overestimation of $T_{\rm g}$. Third, the chemical composition 440 441 of toluene SOA are likely different between Hinks et al. (2017) and Song et al. (2016) because of 442 the differences in the experimental conditions. Specifically, toluene SOA was formed in a Teflon 443 chamber in Hinks et al., while Song et al. used an oxidation flow reactor to generate toluene 444 SOA. The O:C ratios are 0.71 at low RH and 0.63 at high RH based on nano-DESI-HRMS 445 measurements in Hinks et al. (2017), while it was 1.06 based on the aerosol mass spectrometry 446 (AMS) measurements in Song et al. (2016).

447 In addition, different mass loadings may have affected viscosity. Song et al. (2016) measured viscosity at two different mass loadings (60-100 and 600-1000 µg m⁻³) and compared 448 their results to Bateman et al. (2015) (30-50 μ g m⁻³) and Li et al. (2015) (44-125 μ g m⁻³), 449 450 observing little impact of mass loadings on viscosity. We carried out a sensitivity study of mass loadings on viscosity using a set of compounds detected by HRMS. The saturation mass 451 452 concentration was predicted for each component using the molecular corridor approach (Li et al., 453 2016). Assuming that the mass signal intensity is proportional to the total mass concentration of 454 the compound in the mixture, and applying the absorptive partitioning theory (Pankow, 1994), particle-phase concentrations of each compound were predicted to estimate T_{g} at different 455 organic aerosol mass loading values (1-1000 µg m⁻³). The glass transition temperature of the 456 457 SOA mixture decreases as mass loading increases. Viscosity decreases up to two orders of 458 magnitude at low RH, while at high RH they have little difference as shown in Fig. A3. 459 Simultaneous measurements of viscosity and chemical composition with different mass loadings 460 should be performed in future studies.

461

462 **3.3 Biomass Burning Particles**

To further explore the applicability of our viscosity prediction method using elemental composition as measured by HRMS, we performed similar calculations for biomass burning organic particles emitted from test facility burns of subalpine fir and lodgepole pine trees, conducted as a part of the FIREX 2016 campaign (Selimovic et al., 2017). These samples were analyzed by HRMS using two different ionization sources: electrospray ionization (ESI) and atmospheric pressure photoionization (APPI). Mass spectra shown in Fig. 8(a) and (b) indicate that a substantial number of compounds were detected by both methods (109 and 170 470 compounds for subalpine fur and lodgepole pine, respectively). However, pronounced
471 differences are also observed between the ESI and APPI spectra both in terms of the identity and
472 signal intensities of the detected compounds.

473 Glass transition temperatures for the assigned CH and CHO compounds were computed 474 using Eq. (2). Nitrogen and sulfur containing compounds (CHON and CHOS) are not yet 475 covered by Eq. (2) and were therefore excluded from the analysis. CHON and CHOS compounds comprised less than 10% of the detected ion intensity and <15% of the assigned compounds. 476 477 Please note that we do not intend to provide accurate estimates of ambient biomass burning 478 particles (as inorganic components are also not included in this analysis), but we investigate how the use of different ionization methods would lead to variations in our viscosity predictions. T_g of 479 organic mixtures $(T_{g,org})$ were then calculated using the Gordon-Taylor approach with $k_{GT} = 1$, 480 481 assuming that the relative concentration of each compound is proportional to its MS signal intensity. The calculated $T_{g,org}$ values for the mixtures are specified in the legend of Figure 9. For 482 483 both types of mixtures, the calculated $T_{g,org}$ for the APPI MS data is lower than the value 484 calculated based on the ESI MS data with a difference of 32 K for subalpine fir and 11 K for the 485 lodgepole pine. Figure 9 shows the predicted viscosity as a function of RH, assuming D = 10, κ 486 = 0.10 and $k_{\rm GT}$ = 2.5. The difference in $T_{\rm g,org}$ derived from ESI and APPI results in a variation of 487 predicted viscosity at low RH by up to five and two orders of magnitude for subalpine fir and 488 lodgepole pine, respectively.

The difference in the calculated $T_{g,org}$ values is attributed to the chemical profile of the species detected using different ionization techniques as shown in mass spectra in Fig. 8(a) and (b). Van Krevelen diagrams in Fig. 8(c) and (d) illustrate these compositional differences between chemical species detected by ESI and APPI. ESI is more efficient at detection of polar

493 compounds (Kiontke et al., 2016), which typically have higher O:C ratios and therefore would 494 result in higher predicted values of viscosity (Koop et al., 2011; Saukko et al., 2012). APPI 495 enables the detection of nonpolar compounds with lower O:C ratios, in particular polycyclic 496 aromatic hydrocarbons (PAHs), that have low ionization efficiencies when analyzed by ESI MS 497 (Raffaelli and Saba, 2003; Itoh et al., 2006). Due to the complementary nature of these ionization 498 methods, it is most likely that the actual glass transition temperature and viscosity of each type of 499 SOA are somewhere in between the values inferred from ESI and APPI data sets: ESI MS may 500 be viewed as providing the upper limit of viscosity, while APPI MS gives the lower limit. Our 501 results indicate that the use of complementary ionization techniques may help evaluate the 502 associated uncertainty for the prediction of viscosity values based on the elemental composition 503 as measured by HRMS.

504

505 4 Conclusions

506 We have developed a parameterization for calculation of the glass transition temperature of individual SOA compounds with molar mass up to ~ 1100 g mol⁻¹ using the number of carbon, 507 oxygen, and hydrogen atoms. Viscosity of SOA was estimated using the T_{g} -scaled Arrhenius plot 508 509 of viscosity versus T_g/T and the Gordon-Taylor approach to account for mixtures of SOA and 510 water. The fragility parameter D was compiled for organic compounds and we found that D511 approaches a lower limit of ~ 10 (+/- 1.7) as the molar mass increases. The resulting viscosity 512 estimations agree well with measured viscosity of α -pinene and isoprene SOA, validating our 513 method. Using HRMS data, glass transition temperatures of individual components and viscosity 514 of toluene SOA were predicted, also resulting in a good agreement with measurements. 515 However, we note that the predicted viscosities were slightly higher than the measured values

516 suggesting that additional considerations may need to be taken into account. For example, the 517 ionization efficiency of both low and high molar mass compounds may have a pronounced effect 518 on the relative abundance of different classes of compounds in HRMS data. The viscosity 519 prediction method was also applied to biomass burning particles, whose elemental composition 520 was measured using HRMS with two different ionization techniques. Substantial differences in 521 viscosity estimations were obtained using ESI and APPI mass spectra.

522 Figure 10 summarizes the predicted range of viscosity of α -pinene SOA, isoprene SOA, 523 toluene SOA, and biomass burning particles. Isoprene SOA has lower viscosity, reflecting lower 524 glass transition temperature due to relatively low molar mass of isoprene oxidation products. α -525 pinene and toluene SOA have much higher viscosity with a different shape of the RH 526 dependence due to differences in glass transition temperatures and hygroscopicity. Biomass 527 burning particles have moderate viscosity between the two extreme cases. Currently, both 528 predictions and measurements are subject to large uncertainties and variations. Complementary 529 measurements of viscosity and chemical composition employing different ionization techniques are desired to further constrain RH-dependent viscosity in future studies. Current T_{g} 530 531 parameterizations do not consider functionality or molecular structure explicitly and further 532 measurements of $T_{\rm g}$ and viscosity of SOA would allow us to refine the method presented in this 533 study. Nevertheless, current results offer a promising starting point and such simple 534 parameterizations are practical for predicting viscosity of particles as measured by HRMS. The 535 developed viscosity prediction method should also be useful in recent efforts of simulating the 536 distribution of SOA phase state and related properties in regional or global air quality models 537 (e.g., Maclean et al., 2017; Shiraiwa et al., 2017).

539 Appendix A: Conversion of fragility steepness index (m) to fragility (D)

540 Fragility steepness index (*m*) is defined as:

$$m = \lim_{T \to Tg} \frac{d \log \eta}{d(T_g/T)}$$
(A1)

542 Combining Eq. (A1) with Eq. (4) gives:

543
$$m = \lim_{T \to Tg} \frac{d}{d(T_g/T)} \left(-5 + 0.434 \frac{T_0 D}{T - T_0} \right)$$
(A2)

544 Considering that $\eta = 10^{12}$ Pa s at $T = T_g$ (Angell, 1991), and by defining $\Delta x = 1 - T_g/T$, and a

545 combination with Eq. (5) leads to:

$$m = \lim_{\Delta x \to 0} \frac{1}{\Delta x} \left(12 - \left(-5 + 0.434 \frac{\frac{39.17 T_g}{D + 39.17 D}}{\frac{T_g}{1 - \Delta x} - \frac{39.17 T_g}{D + 39.17}} \right) \right)$$
$$= \lim_{\Delta x \to 0} \frac{1}{\Delta x} \left(17 - 0.434 \frac{39.17 T_g D (1 - \Delta x)}{D T_g + 39.17 T_g \Delta x} \right)$$
$$= \lim_{\Delta x \to 0} \frac{(665.89 + 17D)}{(D + 39.17 \Delta x)}$$
$$= \frac{665.89 + 17D}{D}$$
(A3)

Note that Eq. (A3) is derived assuming the high temperature limit of viscosity η_{∞} is equal to 10^{-5} Pa s (Angell, 1991) in the VTF equation (Eq. 3). Similar equations for the relation between *m* and *D* were given by previous studies using different η_{∞} and units (Angell et al., 1994; Angell, 2002; Bones et al., 2012) and applying those gave very similar results in our study.

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- 899

900 **Table 1.** Composition classes and the n_c^0 and b values (K) for glass transition temperature

901 parameterizations obtained by least-squares optimization using the measurements compiled in

Classes	$n_{ m C}^0$	$b_{ m C}$	$b_{ m H}$	$b_{ m CH}$	b_{O}	$b_{ m CO}$
СН	1.96	61.99	-113.33	28.74		
	(±1.81)	(±53.65)	(± 44.47)	(±20.86)		
СНО	12.13	10.95	-41.82	21.61	118.96	-24.38
	(±2.66)	(±13.60)	(±14.78)	(±5.30)	(±9.72)	(±4.21)

902 Koop et al., (2011), Dette et al., (2014) and Rothfuss and Petters (2017).

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906 Table 2. Glass transition temperatures calculated using Eq. (1) and (2) for toluene SOA mixtures

at low relative humidity (low RH < 2%) and high relative humidity (high RH = 75%) conditions.

$T_{\rm g,org}\left({\rm K}\right)$	low RH	high RH
Equation (1)*	299	295
Equation (2)	313	303

908 * Compounds with M > 450 g mol⁻¹ were excluded from the analysis.





910 Figure 1. Characteristic relationships between molecular properties and the glass transition temperature (T_g) of organic compounds. (a) T_g of organic compounds as measured (circles) and 911 estimated with the Boyer-Kauzmann rule (squares) plotted against molar mass. The markers are 912 color-coded by atomic O:C ratio. (b) Measured (circles) and estimated (squares) $T_{\rm g}$ of organic 913 914 compounds plotted against O:C ratio. The markers are color-coded by molar mass. (c) Predicted $T_{\rm g}$ for CHO compounds using a parameterization (Eq. 2) developed in this study compared to 915 measured (circles) and estimated T_g by the Boyer-Kauzmann rule (squares). The solid line shows 916 917 1:1 line and the dashed and dotted lines show 68% confidence and prediction bands, 918 respectively. 919



Figure 2. The Angell plot of viscosity (η) vs. T_g/T . The lines represent different fragility parameter (*D*) values in the range of 5 - 100, with D = 10 (the solid line) used as a base case for this study. A large fragility parameter value is associated with a strong glass former, while fragile materials are associated with lower values. The black dashed line at viscosity of 10^2 Pa s indicates the approximate threshold between liquid and semi-solid states.



927 $^{\text{molar mass (g mol^{-1})}}$ **Figure 3.** Fragility parameter of organic compounds (*D*) plotted against (a) molar mass and (b) 928 atomic O:C ratio. Error bars are standard deviations. The solid red lines represent the fitted 930 curves with fitted equations for (a) D = 602.6/M + 10.3 and (b) D = 14.4 - 2.3(O:C) respectively. 931 Dashed red lines indicate the 95% confidence band.





935 Figure 4. Comparison of measured and predicted viscosity of (a) α -pinene SOA and (b) isoprene 936 SOA at 295 K as a function of RH. The solid lines represent base simulations with the VTF 937 equation, while the dotted line represents viscosity predicted using the WLF equation 938 [parameters: glass transition temperature of dry SOA ($T_{g,org}$), fragility (D), hygroscopicity (κ) and Gordon-Taylor constant (k_{GT})]: (a) 278.5 K, 0.1, 10 and 2.5; (b) 255 K, 0.1, 10 and 2.5. The 939 shaded regions were determined by varying these parameters (a) upper (lower) limit: $T_{g,org} = 300$ 940 941 K (278.5 K), $\kappa = 0.1$ (0.1), D = 20 (10), $k_{\text{GT}} = 2.5$ (2.0); (b) upper (lower limit): $T_{\text{g,org}} = 255$ K (255 K), $\kappa = 0.10$ (0.15), D = 15 (8), $k_{GT} = 2.5$ (4.0). Panel (a): Renbaum-Wolff et al. (2013) data 942 943 represents viscosity for water-soluble portion of SOA; Grayson et al. (2016) data in the panel (a) represents two different mass loadings (121 ug m⁻³; 520 ug m⁻³). Panel (b): The gray box in 944 panel (b) represents estimated viscosity based on bounce measurements of Bateman et al. (2015). 945 946



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Figure 5. Sensitivity calculations for viscosity of α -pinene SOA at 295 K as a function of RH by varying: (a) glass transition temperature of dry SOA ($T_{g,org}$), (b) fragility (D), (c) hygroscopicity

951 (κ), and (d) Gordon-Taylor constant (k_{GT}).





Figure 6. Sensitivity calculations for viscosity of isoprene SOA at 295 K as a function of RH by varying: (a) glass transition temperature of dry SOA ($T_{g,org}$), (b) fragility (*D*), (c) hygroscopicity (κ), and (d) Gordon-Taylor constant (k_{GT}). Data points are measured viscosity by Song et al. (2015) and the gray box represents estimated viscosity based on bounce measurements of Bateman et al. (2015).



961 Figure 7. (a) Molecular corridor of molar mass plotted against volatility of toluene SOA formed under dry conditions (Hinks et al., 2017) color-coded by glass transition temperature (T_{a}) 962 963 estimated using Eq. (2). The upper dashed line indicates the low O:C bound of the molecular 964 corridor (linear alkanes C_nH_{2n+2} with O:C = 0), and the lower dotted line indicates the high O:C 965 bound (sugar alcohols $C_nH_{2n+2}O_n$ with O:C = 1). (b) Comparison of measured (markers) and modeled (lines) viscosity of toluene SOA at 295 K as a function of RH. Viscosities were 966 967 calculated using fragility (D) of 13, the hygroscopicity (κ) of 0.25 and the Gordon-Taylor constant (k_{GT}) of 3.0 with different glass transition temperatures of dry SOA ($T_{g,org}$) as estimated 968 using Eq. (1) or (2) under low and high RH conditions. The shaded regions were calculated by 969 varying those parameters: $T_{g,org} = 313$ K (295 K), $\kappa = 0.20$ (0.25), D = 13 (10), $k_{GT} = 2.5$ (3.5) for 970 the upper (lower) limit. Mass loadings were 23 µg m⁻³ for LRH and 8 µg m⁻³ for HRH (Hinks et 971 972 al., 2017).



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Figure 8. Mass spectra of biomass burning organic particles collected from test burns of (a) subalpine fir and (b) lodgepole pine as measured by high resolution mass spectrometry with two ionization techniques: electron spray ionization (ESI, red) and atmospheric pressure photoionization (APPI; blue). Numbers of elemental formulas identified by ESI (red), APPI (blue) and both modes (black) are also specified. Van Krevelen plots of the compounds identified by ESI (red) and APPI (blue) mode in BBOA from burning of (c) subalpine fir and (d) lodgepole pine.



Figure 9. Predicted viscosity for biomass burning particles of (a) subalpine fir and (b) lodgepole pine trees as measured by high resolution mass spectrometry with two ionization techniques: electrospray ionization (ESI, red) and atmospheric pressure photoionization (APPI; blue). $T_{g,org}$ are specified in the figure legend and other used parameters are fixed to $\kappa = 0.1$, D = 10, $k_{GT} = 2.5$.

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Figure 10. Summary of predicted range of viscosity of α-pinene SOA (red), isoprene SOA

- 995 (blue), toluene SOA (purple), and biomass burning particles (green).
- 996



Figure A1. (a) T_g of organic compounds as measured (circles) and estimated with the Boyer-Kauzmann rule (squares) plotted against the inverse molar mass. The markers are color-coded by atomic O:C ratio. (b) Predicted T_g for CHO compounds using a parameterization (Eq. 2) developed in this study compared to measured T_g (circles). The solid line shows 1:1 line and the dashed and dotted lines show 68% confidence and prediction bands, respectively.



1004 Measured $T_g[K]$ Measured $T_g[K]$ 1005 **Figure A2.** (a) Comparison of measured and estimated T_g by the Boyer-Kauzmann rule for 251 1006 organic compounds (Koop et al., 2011; Dette et al., 2014; Rothfuss and Petters 2017) with their 1007 measured T_m available. The markers are color-coded by molar mass. (b, c) Predicted T_g using Eq. 1008 (2) compared with (b) measured T_g for CH compounds and (c) predicted T_g using Eq. (1) for 1009 CHO compounds. The solid line shows 1:1 line. Solid circle markers represent organic 1010 compounds as compiled in Koop et al. (2011) and open square marker represent SOA oxidation 1011 products in Shiraiwa et al. (2014) in panel (c).



Figure A3. Effect of mass loading on predicted viscosity for toluene SOA. Solid lines represent the predicted viscosity with Eq. (2) using chemical composition of toluene SOA formed at low

RH. Viscosity was predicted with different mass loadings ranging from 1-1000 µg m⁻³. Markers

and shaded boxes represent experimentally measured viscosity values. Song et al. (2016) mass loadings were 60-100 and 600-1000 μ g m⁻³. Bateman et al., (2015) and Li et al., (2015) mass

loadings were 30-50 µg m⁻³ and 44-125 µg m⁻³, respectively.