



1 **Comparison of secondary organic aerosol formation from**
2 **toluene on initially wet and dry ammonium sulfate particles**

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19 **Abstract**

20 The formation of secondary organic aerosol (SOA) has been widely studied in the
21 presence of dry seed particles at low relative humidity (RH). At higher RH, seed
22 particles can exist as dry or wet particles. Here, we investigated the formation of SOA
23 from the photooxidation of toluene using an oxidation flow reactor under a range of
24 OH exposures on initially wet or dry ammonium sulfate (AS) seed particles at an RH
25 of 68%. At an OH exposure of 4.66×10^{10} molecules cm^{-3} s, the ratio of the SOA yield
26 on wet AS seeds to that on dry AS seeds was 1.31 ± 0.02 . However, this ratio decreased
27 to 1.01 ± 0.01 at an OH exposure of 5.28×10^{11} molecules cm^{-3} s. The decrease in the
28 ratios of SOA yields as the increase of OH exposure may be due to the early
29 deliquescence of initially dry AS seeds after coated by highly oxidized toluene-derived
30 SOA. SOA formation lowered the deliquescence RH of AS and resulted in the uptake
31 of water by both AS and SOA. Hence the initially dry AS seeds contained aerosol liquid
32 water (ALW) soon after a large fraction of SOA formed and the SOA yield and ALW
33 approached those of the initially wet AS seeds as OH exposure and ALW increased.
34 However, a higher oxidation state of the SOA on initially wet AS seeds than that on dry
35 AS seeds was observed at all levels of OH exposure. The difference in mass fractions
36 of m/z 29, 43 and 44 of SOA mass spectra indicated that SOA formed on initially wet
37 seeds may be enriched in earlier-generation products containing carbonyl functional
38 groups at low OH exposures and later-generation products containing acidic functional
39 groups at high exposures. Our results suggest that AS dry seeds soon turn to at least
40 partially deliquesced particles during SOA formation and more studies on the interplay



41 of SOA formation and ALW are warranted.



42 **1. Introduction**

43 Secondary organic aerosol (SOA) is an important component of atmospheric particulate
44 matter, which influences air quality, climate and human health (Hallquist et al., 2009).
45 SOA is mainly formed via the oxidation of volatile organic compounds (VOCs),
46 followed by partitioning to the condensed phase. Traditional atmospheric chemical
47 transport models largely underestimate the levels of SOA (de Gouw et al., 2005;
48 Volkamer et al., 2006; Hodzic et al., 2010) and the degree of oxidation (Rudich et al.,
49 2007; Ng et al., 2010). SOA yields in traditional atmospheric chemical transport models
50 are obtained from smog chamber experiments using dry seed particles (Barsanti et al.,
51 2013; Mahmud and Barsanti, 2013) under dry conditions. Yet, atmospheric relative
52 humidity is often sufficiently high that aerosols often exist as wet aerosols, containing
53 a large amount of aerosol liquid water (ALW) (Liao and Seinfeld, 2005; Lee and Adams,
54 2010; Guo et al., 2015; Nguyen et al., 2016). The presence of ALW in wet aerosols may
55 enhance SOA formation by facilitating the partitioning of semivolatile organic
56 compounds and the uptake of water-soluble gases through aqueous-phase reactions
57 (Hennigan et al., 2008; Lim et al., 2010; Ervens et al., 2011; Lee et al., 2011; Sareen et
58 al., 2017). ALW may also promote photodegradation of dissolved SOA (Romonosky et
59 al., 2014). Therefore, SOA formation under atmospherically relevant relative humidity
60 needs to be better constrained in atmospheric chemical transport models.

61 Aromatic hydrocarbons constitute a large fraction of the total non-methane
62 hydrocarbons in the urban atmosphere (Calvert et al., 2002) and account for a
63 significant fraction of SOA in urban areas (Ding et al., 2012; Zhao et al., 2017). Toluene



64 is the most abundant aromatic hydrocarbon (Calvert et al., 2002; Zhang et al., 2016)
65 and SOA yields from the photooxidation of toluene on dry or wet ammonium sulfate
66 (AS) seeds has been studied by varying the RH in smog chambers. Kamens et al. (2011)
67 observed higher yields of SOA from toluene at higher RHs. They attributed this increase
68 to the initially wet seed particles. On the other hand, Edney et al. (2000) reported that
69 wet seeds had no effect on the SOA yields of toluene compared with dry seeds. In these
70 studies, different RHs used for dry and wet seeds experiments may influence the gas-
71 phase chemistry and complicate the comparison of SOA formation.

72 SOA formation on initially dry and wet AS seeds has been compared using
73 oxidation flow reactors at same RHs (Wong et al., 2015; Faust et al., 2017). Faust et al.
74 (2017) found a 19% enhancement in the SOA yield of toluene on wet AS seeds over
75 that on dry AS seeds at 70% RH. However, at such high RH, the initially dry and water-
76 free AS seed particles can uptake water upon SOA formation because SOA themselves
77 can be hygroscopic and they can also lower the deliquescence RH of the AS seeds
78 (Takahama et al., 2007; Smith et al., 2011, 2012, 2013). The potential influence of SOA
79 formation on the physical state of the initially dry seeds as well as and the overall water
80 uptake by the aged particles was not explicitly discussed. In addition, the hydroxyl
81 radicals (OH) exposure in their study was approximately 2×10^{11} molecules cm^{-3} s,
82 equivalent to about 1.5 days of oxidation in the atmosphere assuming an ambient OH
83 concentration of 1.5×10^6 molecules cm^{-3} (Mao et al., 2009). Atmospheric particles can
84 undergo oxidation for as long as 1-2 weeks (Balkanski et al., 1993).

85 In this study, SOA formation from the photooxidation of toluene was investigated



86 in an oxidation flow reactor at an RH of 68% under a wide range of OH exposures using
87 initially wet or dry AS seed particles. The yields and composition of SOA as well as the
88 estimated ALW contents for the initially wet and dry seeds are compared. We found
89 that as OH exposure increased, the SOA yield and ALW of the initially dry seeds
90 approached those of the initially wet seeds while the wet seeds yielded SOA of a higher
91 degree of oxidation than the dry seeds did at all exposure levels.

92 **2. Materials and methods**

93 **2.1 Generation of seed particles**

94 A schematic of the experimental setup is shown in Fig. 1. AS seed particles were
95 generated from an aqueous AS solution (Sigma-Aldrich) using an atomizer (TSI 3076,
96 TSI Inc., USA). In experiments using dry seeds, the atomized aqueous AS droplets
97 passed through a silica gel diffusion dryer so that the RH was reduced to less than 30%
98 at which AS effloresced, while in experiments using wet seeds, they bypassed the
99 diffusion dryer. The dry or wet seed particles then entered and mixed with a humidified
100 N₂/O₂/O₃ flow in an oxidation flow reactor. The RH in the flow reactor was at 68%,
101 which lies between the efflorescence and deliquescence RH of AS (Seinfeld and Pandis,
102 2006), so that the seed particles remained in their original phase with the wet particles
103 containing ~18.6 μg m⁻³ ALW and the dry particles anhydrous before reaction started.
104 Hereafter, the experiments using initially wet and dry AS seed particles are simplified
105 as wet and dry AS seeds, respectively. “Wet” and “dry” refer to the initial state of the
106 seed particles before SOA formation.

107 When atomizing a given AS solution, the diameter of wet AS droplets is much



108 larger than that of dry AS particles due to the water uptake of AS (Chan et al., 1992),
109 resulting in a larger surface area of seed particles. Previous studies have demonstrated
110 that a large surface area of seed particles may increase the SOA yields by reducing the
111 wall loss of organic vapors (Matsunaga and Ziemman, 2010, Zhang et al., 2014, 2015;
112 Huang et al., 2016; Krechmer et al., 2016). To obtain seed particles of comparable
113 surface areas, we atomized 0.013 mM and 0.015 mM of the AS solution for wet and
114 dry AS seeds, respectively. As shown in Fig. S1, the surface area distribution of wet AS
115 seeds was similar to that of dry AS seeds. Because of the difference in AS concentration
116 between the stock solutions used, wet AS seeds had a mean diameter of 88 nm and were
117 slightly smaller than dry AS seeds which had a mean diameter of 102 nm. The total
118 surface area of wet AS seeds was 21% larger than that of dry AS seeds. The mass
119 loading of wet and dry AS seeds was 31.0 and 24.2 $\mu\text{g m}^{-3}$, respectively.

120 **2.2 Oxidation flow reactor**

121 SOA formation from the photooxidation of toluene on initially dry or wet seeds was
122 investigated in a potential aerosol mass (PAM) oxidation flow reactor, which has been
123 described in detail elsewhere (Kang et al., 2007, 2011; Lambe et al., 2011a, 2015; Liu
124 et al., 2017). Briefly, a PAM chamber is a continuous oxidation flow reactor using high
125 and controlled levels of oxidants to oxidize gaseous precursors to produce SOA. The
126 chamber used in this study had a volume of approximately 19 L (length 60 cm, diameter
127 20 cm). The total flow rate in the PAM chamber was set at 3 L min^{-1} using mass flow
128 controllers, resulting in a residence time of approximately 380 s. The RH and
129 temperature of the PAM outflow were measured continuously (HMP 110, Vaisala Inc,



130 Finland) and stabilized at approximately 68% and 20 °C, respectively. High OH
131 exposures were realized through the photolysis of ozone irradiated by a UV lamp ($\lambda =$
132 254 nm) in the presence of water vapor. Ozone was produced by an ozone generator
133 (1000BT-12, ENALY, Japan) via the irradiation of pure O₂. The OH concentration was
134 adjusted by varying the concentration of ozone in the PAM chamber from 0.4 ppm to
135 4.3 ppm. The corresponding upper limit of OH exposure at these operating conditions
136 ranged from 4.66×10^{10} molecules cm⁻³ s to 5.28×10^{11} molecules cm⁻³ s, equivalent to
137 0.36 to 4.08 days of atmospheric oxidation assuming an ambient OH concentration of
138 1.5×10^6 molecules cm⁻³ (Mao et al., 2009). The addition of toluene may reduce the OH
139 exposure. The upper limit of OH exposure was determined by measuring the decay of
140 SO₂ (Model T100, TAPI Inc., USA) in the absence of toluene, following procedures
141 described elsewhere (Kang et al., 2007; Lambe et al., 2011a). Peng et al. (2016) found
142 that non-OH chemistry, including photolysis at $\lambda = 254$ nm and reactions with O(¹D),
143 O(³P) and O₃, may play an important role in oxidation flow reactors. In this study, the
144 PAM reactor was operated at water vapor mixing ratios above 0.5% and external OH
145 reactivity below 20 s⁻¹. Non-OH chemistry is expected to play a negligible role under
146 these conditions (Peng et al., 2016).

147 Before and after each experiment, the PAM reactor was cleaned under an OH
148 exposure of $\sim 1 \times 10^{12}$ molecules cm⁻³ s until the mass concentration of background
149 particles dropped below 3 $\mu\text{g m}^{-3}$. After characterizing dry or wet AS seed particles for
150 half an hour, the UV lamp was turned on to oxidize the background gases at five
151 different OH levels to measure the concentrations of background organics. A toluene



152 mixture (29.6 ppm in nitrogen) with a flow rate of 0.013 L min^{-1} was then introduced
153 to initiate SOA formation. The initial concentration of toluene in the PAM reactor was
154 approximately 138 ppb. SOA was measured for at least an hour at each of the five OH
155 levels.

156 **2.3 Characterization of non-refractory components**

157 The AS/SOA mixed particles were characterized for the chemical composition of non-
158 refractory components including organics, sulfate and ammonium as well as the
159 elemental ratios of organics using a high-resolution time-of-flight aerosol mass
160 spectrometer (hereafter AMS, Aerodyne Research Incorporated, USA) (DeCarlo et al.,
161 2006). The instrument was operated in the high sensitivity V-mode and the high
162 resolution W-mode alternating every one minute. The toolkit Squirrel 1.57I and Pika
163 1.16I were used to analyze the AMS data. The molar ratios of hydrogen to carbon (H:C)
164 and oxygen to carbon (O:C) were determined using the Aiken method (Aiken et al.,
165 2007, 2008). The ionization efficiency of the AMS was calibrated using 300 nm
166 ammonium nitrate particles. The particle-free matrix air, obtained by passing the air
167 flow from the PAM reactor through a HEPA filter, was measured for at least 20 min
168 before each experiment to determine the signals from major gases.

169 The collection efficiency (CE) of an AMS is dependent on the chemical
170 composition and acidity as well as the phase state of particles (Matthew et al., 2008;
171 Middlebrook et al., 2012). Matthew et al. (2008) found that the CE for solid particles
172 thickly coated with liquid organics was 100%. In this study, experiments were
173 conducted at an RH of 68%, exceeding the RH threshold for the semisolid-to-liquid



174 phase transition for toluene-derived SOA (Bateman et al., 2015; Song et al., 2016). A
175 CE of 1 was used for processing all AMS data since the concentration of sulfate
176 measured with the AMS varied by less than 5% of the average mass of sulfate after
177 coated by SOA for both wet and dry AS seeds conditions. For the quantification of
178 SOA, the contribution from background organic aerosols was subtracted from the total
179 organic aerosols. The ratio of SOA mass to background organic mass ranged from 7 to
180 59, indicating that the contribution from background organics was negligible. Aerosol
181 particles typically pass through a silica gel diffusion dryer to remove ALW before they
182 are measured by AMS. However, this may lead to some losses of semivolatile organics
183 through reversible partitioning (Wong et al., 2015; Faust et al., 2016). In this study, the
184 AS/SOA mixed particles stream passed through and bypassed a diffusion dryer
185 alternately before they were measured by AMS. Overall less than 8% of SOA were lost
186 for wet and dry AS seeds after passing the diffusion dryer (Fig. S2), possibly due to
187 reversible partitioning of the SVOCs. In this paper, the data reported are those
188 bypassing the diffusion dryer.

189 A scanning mobility particle sizer (SMPS, TSI Incorporated, USA, classifier model
190 3082, CPC model 3775) was used to measure particle number concentrations and size
191 distributions. Particle size ranged from 15 nm to 661 nm.

192 To evaluate the influence of seed surface area on SOA formation, we conducted
193 another experiment at OH exposure of 4.66×10^{10} molecules cm^{-3} s with 50% of the seed
194 surface area used in the wet AS experiment. The difference in SOA concentration was
195 approximately 1% between these two experiments. Hence the 20% difference in seed



196 surface area as well as the difference in mass loadings between wet and dry AS particles
197 cannot account for the difference in SOA yield to be discussed below.

198 **2.4 Estimation of aerosol liquid water (ALW) content**

199 The ALW content of the initially dry AS was zero. However, as reactions proceed, SOA
200 themselves can uptake water and also lower the deliquescence RH of AS, leading to
201 water uptake by AS and some fractions of AS in aqueous phase. The ALW contents of
202 AS (ALW_{AS}) and toluene-derived SOA (ALW_{SOA}) were estimated from the following
203 equations (Kreidenweis et al., 2008):

$$204 \quad ALW_{AS} = V_{AS} \kappa_{AS} f \frac{\alpha_w}{1 - \alpha_w} \rho_w \quad (1)$$

$$205 \quad ALW_{SOA} = V_{SOA} \kappa_{SOA} \frac{\alpha_w}{1 - \alpha_w} \rho_w \quad (2)$$

206 where V_{AS} and V_{SOA} represent the volume concentrations of dry AS and SOA particles,
207 κ_{AS} is the hygroscopicity parameter of AS particles obtained from Kreidenweis et al.
208 (2008), κ_{SOA} is the hygroscopicity parameter of toluene-derived SOA calculated using
209 the linear correlation between κ_{SOA} and the O:C ratios of SOA proposed by Lambe et al.
210 (2011b), the term f is the fraction of AS particles that dissolved, α_w is the water activity
211 and ρ_w is the density of water (1.0 g cm^{-3}). Here, α_w was assumed to be equivalent to
212 RH/100 for simplicity. The volume concentrations of dry AS and SOA particles were
213 estimated from the measured mass concentration of AS and SOA assuming their
214 respective particle densities to be 1.77 g cm^{-3} and 1.4 g cm^{-3} (Ng et al., 2007).

215 For the initially wet AS seeds, all AS particles were completely aqueous and
216 therefore $f = 1$. For the initially dry AS seeds, before reactions, the AS particles were



217 completely dry and $f = 0$. After reactions, the AS particles became partially or entirely
218 deliquesced upon the formation of toluene-derived SOA. The dissolved fraction of AS
219 particles was regulated by the liquidus curve of the deliquescence relative humidity
220 (DRH(ε)) of AS particles coated with toluene-derived SOA (Smith et al., 2013):

$$221 \quad f = \begin{cases} \frac{\varepsilon(1-\varepsilon_D)}{\varepsilon_D(1-\varepsilon)} & \text{for } \varepsilon < \varepsilon_D \\ 1 & \text{for } \varepsilon \geq \varepsilon_D \end{cases} \quad (3)$$

222 The term ε is the volume fraction of SOA. The term ε_D , representing the volume fraction
223 of organics at which the mixture of SOA and AS particles deliquesced at an RH of 68%,
224 was estimated to be 0.75 based on the liquidus curve.

225 **3. Results and discussion**

226 **3.1 SOA yields**

227 Figure 2a shows SOA yields from the photooxidation of toluene on initially wet and
228 dry AS seed particles as a function of OH exposure. The SOA yield was calculated as
229 the SOA mass divided by the mass of reacted toluene. The mass of reacted toluene was
230 calculated from the OH exposure and the rate constant of the reaction between toluene
231 and OH (Atkinson and Arey, 2003). The uncertainty in the SOA yields fully reflected
232 the uncertainty in the calculation of the SOA mass. In both cases, SOA yields first
233 exhibited an increase, followed by a decrease as the level of OH exposure increased.
234 This trend may be due to the transition of functionalization reactions to fragmentation
235 ones (Kroll et al., 2009; Lambe et al., 2011a). The SOA yields for dry and wet AS seeds
236 were 0.18–0.31 and 0.22–0.36, respectively, significantly higher than the value of
237 0.0059 observed in an oxidation flow reactor under comparable conditions (Faust et al.,



238 2017) and the value of 0.09 obtained in another PAM chamber at 30% RH in the
239 absence of seed particles (Kang et al., 2007). Faust et al. (2017) attributed their
240 significantly lower yields than typical literature values of 0.09–0.30 (Lambe et al.,
241 2011a; Ng et al., 2007) to the wall loss of particles and the fragmentation of organics in
242 their flow reactor. On the other hand, the SOA yields we obtained are lower than 0.30–
243 0.37 from smog chamber experiments conducted at a similar temperature, SOA mass
244 loading and OH exposure but a lower RH with dry AS seeds (Ng et al., 2007;
245 Hildebrandt et al., 2009). Note that the wall loss of particles was not corrected in this
246 study, so the SOA yields may be underestimated. As wet and dry AS seeds in this study
247 had similar particle number size distributions, the wall loss of particles would not affect
248 the comparison of SOA yield between wet and dry AS seeds.

249 As shown in Fig. 2a, a higher SOA yield was observed for wet AS seeds than for
250 dry AS seeds at the same OH exposure and the difference in SOA yield decreased as
251 the OH exposure increased. The ratio of SOA yields on wet AS seeds to those on dry
252 AS seeds was 1.31 ± 0.02 at an OH exposure of 4.66×10^{10} molecules cm^{-3} s but
253 decreased to 1.01 ± 0.01 when the OH exposure was increased to 5.28×10^{11} molecules
254 cm^{-3} s (Fig. 2b). These ratios are comparable to the 1.19 ± 0.05 observed by Faust et al.
255 (2017) at an OH exposure of approximately 2.0×10^{11} molecules cm^{-3} s.

256 The formation of SOA on initially dry AS particles may alter the deliquescence
257 relative humidity (DRH) of AS particles. Smith et al. (2013) found that when coated
258 with toluene-derived SOA, the DRH of AS particles decreased from 80% to 58% as the
259 organic volume fraction increased from 0 to 0.8. Therefore, coating AS particles with



260 toluene-derived SOA can change the physical state of initially dry AS seeds and
261 increase the content of $ALW_{AS, dry}$. As shown in Fig. 3a, after reactions, the mass
262 concentrations of ALW_{tot} ($= ALW_{SOA} + ALW_{AS}$) and ALW_{SOA} increased for both wet
263 and dry seeds as the OH exposure increased. The uncertainties for ALW_{SOA} and ALW_{AS}
264 were 22% and less than 3%, respectively. They reflect the uncertainties in κ and volume
265 concentrations of AS and SOA. The increase in $ALW_{tot, wet}$ was due to the increase in
266 $ALW_{SOA, wet}$ while the increase in $ALW_{tot, dry}$ was driven by the increase in $ALW_{AS, dry}$ at
267 lower OH exposure and by $ALW_{SOA, dry}$ at higher OH exposures. At OH exposure of
268 4.66×10^{10} molecules cm^{-3} s, $ALW_{AS, dry}$ increased from 0 to $6.2 \mu g m^{-3}$ after reactions
269 due to the partial deliquescence ($f=0.43$) of the originally dry AS particles after SOA
270 formation. The difference in $ALW_{AS, dry}$ and $ALW_{AS, wet}$ narrowed and the ALW_{total} of
271 initially dry AS seeds partially resembled those of the wet ones. At OH exposure
272 between 1.66×10^{11} and 5.28×10^{11} molecules cm^{-3} s, the total final organic volume
273 fraction increased to approximately 0.8 and the initially dry AS particles entirely
274 deliquesced after reactions. Based on the reported SOA yield, initial toluene
275 concentration, OH exposure and assumed concentrations of AS seeds ($\sim 10\text{--}40 \mu g m^{-3}$)
276 in Faust et al. (2017), we estimated that an upper limit of 48% of the initially dry AS
277 seeds has deliquesced in their study. Similar to this study, SOA coatings on seed
278 particles may change the physical state of initially dry seeds and lower the difference
279 of SOA yields between initially dry and wet seeds experiments.

280 3.2 Chemical composition of SOA



281 Figure 4 shows the high-resolution mass spectra of SOA for initially wet and dry AS
282 seeds at OH exposures of 4.66×10^{10} molecules cm^{-3} s and 5.28×10^{11} molecules cm^{-3} s.
283 For both types of AS seeds, at an OH exposure of 4.66×10^{10} molecules cm^{-3} s, the most
284 prominent peaks were m/z 29 and 43, followed by m/z 28 and 44. m/z 29 was dominated
285 by ion CHO^+ , a tracer for alcohols and aldehydes (Lee et al., 2012). The m/z 28 and m/z
286 44 signals, respectively dominated by CO^+ and CO_2^+ , are tracers for organic acids (Ng
287 et al., 2010). At the OH exposure of 5.28×10^{11} molecules cm^{-3} s, the dominant peaks
288 were m/z 28 and 44, followed by m/z 29 and 43. The increase of mass fractions of the
289 oxygen-containing ions in the SOA mass spectra at a relatively high OH exposure
290 suggests the formation of more oxidized organic aerosols. Furthermore, ions with $m/z >$
291 80 accounted for a negligible fraction of total SOA, suggesting that oligomerization
292 might not be important in these experiments. On the basis of the mass fraction of ions,
293 Fig. S3 shows that as OH exposure increased, the difference (wet minus dry) in the
294 spectra of toluene-derived SOA changed from positive in m/z 29 (CHO^+) and m/z 43
295 ($\text{C}_2\text{H}_3\text{O}^+$) to m/z 28 (CO^+) and m/z 44 (CO_2^+). The increase in OH exposure resulted in
296 a change from more alcohols or aldehydes to more organic acids in the wet seeded case
297 when compared to the dry seeded case.

298 Fragments derived from the AMS data have been extensively used to infer the bulk
299 compositions and evolution of organic aerosols (Zhang et al., 2005; Ng et al., 2010;
300 Heald et al., 2010). Here we used the approach of Ng et al. (2010) and plotted the
301 fractions of the total organic signal at m/z 43 (f_{43}) vs. m/z 44 (f_{44}) as well as the triangle
302 based on the analysis of ambient AMS data (Fig. 5). Ng et al. (2010) proposed that



303 aging would cause f_{43} and f_{44} to converge toward the triangle apex ($f_{43} = 0.02$, $f_{44} = 0.30$).
304 For both wet and dry AS seeds, f_{43} first increased and then decreased with the increase
305 of OH exposure, while f_{44} increased all the time. This reversing trend of f_{43} was the
306 result of the increase and subsequent decrease in $C_2H_3O^+$ (Fig. S4), an indicator of
307 products containing carbonyl functional groups. It was also observed for SOA formed
308 from other precursors such as alkanes and naphthalene (Lambe et al., 2011b). Before
309 the decrease in f_{43} , SOA formed on wet AS seeds had higher f_{43} and similar f_{44} to SOA
310 formed on dry AS seeds at the same OH exposure. As OH exposure increased, SOA
311 formed on wet AS seeds had higher f_{44} and lower f_{43} than SOA formed on dry AS seeds.
312 The f_{43} - f_{44} plot supports our earlier assertion that as OH exposure increased, the reaction
313 products changed from earlier-generation products containing carbonyl functional
314 groups to later-generation products containing acidic functional groups. In addition, as
315 OH exposure increased, SOA formed on wet AS seeds initially had more earlier-
316 generation products but later had more acidic later-generation products than SOA
317 formed on dry AS seeds, likely due to the enhanced partitioning of these products on
318 initially wet AS seeds and/or enhanced uptake of water-soluble gases through aqueous
319 phase reactions.

320 Figure 6 shows the changes in H:C and O:C ratios as a function of OH exposure in
321 a Van Krevelen diagram (Heald et al., 2010). The standard deviations for H:C and O:C
322 values were both less than 0.01. The O:C ratios for dry and wet AS seeds were in the
323 ranges of 0.59–0.89 and 0.63–0.95, respectively. At the same OH exposure, SOA on
324 wet AS seeds had both higher O:C ratios and estimated average carbon oxidation state



325 (OS_C) ($OS_C \approx 2 \times O:C - H:C$) (Kroll et al., 2011) than dry AS seeds had. Fig. 6 also
326 shows some of the identified SOA products from the photooxidation of toluene (Bloss et
327 al., 2005; Hamilton et al., 2005; Sato et al., 2007). The elevated OS_C (exceeding 0.5)
328 could only be due to the formation of highly oxygenated small acids such as pyruvic
329 acid ($OS_C = 0.67$), glycolic acid ($OS_C = 1$), formic acid ($OS_C = 2$), oxalic acid ($OS_C =$
330 3), malonic acid ($OS_C = 1.33$) and glyoxylic acid ($OS_C = 2$). Small acids may be
331 important products of toluene-derived SOA at high OH exposures. Fisseha et al. (2004)
332 found that small organic acids accounted for 20–45% of SOA from the photooxidation
333 of 1,3,5-trimethylbenzene. The higher OS_C at high OH exposures for wet AS seeds
334 might suggest that these small acids were more abundant, likely due to their enhanced
335 retention in the presence of ALW and/or the more efficient uptake of OH radicals by
336 wet AS seeds and further oxidation reactions in aqueous phase (Ruehl et al., 2013). The
337 change in the slope of H:C vs O:C is consistent with the earlier analysis that the
338 mechanism of SOA formation changed from functionalization dominated by the
339 addition of alcohol/peroxide (Heald et al., 2010; Ng et al., 2011) at low exposures to
340 the addition of both acid and alcohol/peroxide functional groups without fragmentation,
341 and/or the addition of acid groups with fragmentation at high exposures.

342 **3.3 Atmospheric implications**

343 In this work, yields and composition of SOA formed from the photooxidation of toluene
344 on initially wet and dry AS seeds were compared over a wide range of OH exposures,
345 covering the transition from functionalization reactions to fragmentation reactions. We
346 found that the ratio of SOA yield on wet AS seeds to that on dry AS seeds decreased



347 from 1.31 to 1.01 as the OH exposure increased from 4.66×10^{10} to 5.28×10^{11} molecules
348 $\text{cm}^{-3} \text{ s}$. This decrease coincides with the decrease of differences in ALW between the
349 wet and dry cases, which may be due to water uptake by SOA as well as the early
350 deliquescence of dry AS particles as a result of SOA formation.

351 In addition to relatively higher SOA yields, higher O:C and OS_c of SOA derived
352 from the photooxidation of toluene were also observed on initially wet AS seeds.
353 Particularly, the O:C in the presence of initially wet AS seeds could be as high as 0.95.
354 Chen et al. (2015) observed large gaps between laboratory and ambient measured O:C
355 of OA and suggested that OA having a high O:C (> 0.6) was required to bridge these
356 gaps. The multiphase oxidation of toluene in the presence of wet aerosols may be a
357 pathway to contribute to this gap. However, the relative importance of such chemistry
358 to the evolution of ambient OA remains unclear.

359 Our results suggest that dry seeds would quickly turn to at least partially
360 deliquesced particles upon SOA formation under moderate RH conditions. Since
361 ambient RH is rarely at such low values that inorganic particles remain dry even after
362 SOA formation, more laboratory and field studies are needed to elucidate the formation
363 and evolution of OA in wet aerosols.



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371 **References**

- 372 Aiken, A. C., DeCarlo, P. F., and Jimenez, J. L.: Elemental Analysis of Organic Species with
373 Electron Ionization High-Resolution Mass Spectrometry, *Anal. Chem.*, 79, 8350-
374 8358, <https://doi.org/10.1021/ac071150w>, 2007.
- 375 Aiken, A. C., DeCarlo, P. F., Kroll, J. H., Worsnop, D. R., Huffman, J. A., Docherty, K. S.,
376 Ulbrich, I. M., Mohr, C., Kimmel, J. R., Sueper, D., Sun, Y., Zhang, Q., Trimborn,
377 A., Northway, M., Ziemann, P. J., Canagaratna, M. R., Onasch, T. B., Alfarra, M. R.,
378 Prevot, A. S. H., Dommen, J., Duplissy, J., Metzger, A., Baltensperger, U., and
379 Jimenez, J. L.: O/C and OM/OC Ratios of Primary, Secondary, and Ambient Organic
380 Aerosols with High-Resolution Time-of-Flight Aerosol Mass Spectrometry, *Environ.*
381 *Sci. Technol.*, 42, 4478-4485, <https://doi.org/10.1021/es703009q>, 2008.
- 382 Atkinson, R., and Arey, J.: Atmospheric Degradation of Volatile Organic Compounds, *Chem.*
383 *Rev.*, 103, 4605-4638, <https://doi.org/10.1021/cr0206420>, 2003.
- 384 Balkanski, Y. J., Jacob, D. J., Gardner, G. M., Graustein, W. C., and Turekian, K. K.:
385 Transport and residence times of tropospheric aerosols inferred from a global three-
386 dimensional simulation of 210Pb, *J. Geophys. Res.-Atmos.*, 98, 20573-20586,
387 <https://doi.org/10.1029/93JD02456>, 1993.
- 388 Barsanti, K. C., Carlton, A. G., and Chung, S. H.: Analyzing experimental data and model
389 parameters: implications for predictions of SOA using chemical transport models,
390 *Atmos. Chem. Phys.*, 13, 12073-12088, <https://doi.org/10.5194/acp-13-12073-2013>,
391 2013.
- 392 Bateman, A. P., Bertram, A. K., and Martin, S. T.: Hygroscopic Influence on the Semisolid-
393 to-Liquid Transition of Secondary Organic Materials, *J. Phys. Chem. A*, 119, 4386-
394 4395, <https://doi.org/10.1021/jp508521c>, 2015.
- 395 Bloss, C., Wagner, V., Jenkin, M. E., Volkamer, R., Bloss, W. J., Lee, J. D., Heard, D. E.,
396 Wirtz, K., Martin-Reviejo, M., Rea, G., Wenger, J. C., and Pilling, M. J.:
397 Development of a detailed chemical mechanism (MCMv3.1) for the atmospheric
398 oxidation of aromatic hydrocarbons, *Atmos. Chem. Phys.*, 5, 641-664,
399 <https://doi.org/10.5194/acp-5-641-2005>, 2005.
- 400 Calvert, J. G., Atkinson, R., Becker, K. H., Kamens, R. M., Seinfeld, J. H., Wallington, T. H.,
401 and Yarwood, G.: *The Mechanisms of Atmospheric Oxidation of the Aromatic*
402 *Hydrocarbons*, Oxford University Press, New York, 556 pp., 2002.
- 403 Chen, Q., Heald, C. L., Jimenez, J. L., Canagaratna, M. R., Zhang, Q., He, L.-Y., Huang, X.-
404 F., Campuzano-Jost, P., Palm, B. B., Poulain, L., Kuwata, M., Martin, S. T., Abbatt,
405 J. P. D., Lee, A. K. Y., and Liggio, J.: Elemental composition of organic aerosol: The
406 gap between ambient and laboratory measurements, *Geophys. Res. Lett.*, 42,
407 2015GL063693, <https://doi.org/10.1002/2015GL063693>, 2015.
- 408 de Gouw, J. A., Middlebrook, A. M., Warneke, C., Goldan, P. D., Kuster, W. C., Roberts, J.
409 M., Fehsenfeld, F. C., Worsnop, D. R., Canagaratna, M. R., Pszenny, A. A. P.,
410 Keene, W. C., Marchewka, M., Bertman, S. B., and Bates, T. S.: Budget of organic
411 carbon in a polluted atmosphere: Results from the New England Air Quality Study in
412 2002, *J. Geophys. Res.-Atmos.*, 110, D16305, <https://doi.org/10.1029/2004JD005623>,
413 2005.



- 414 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
415 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J.
416 L.: Field-Deployable, High-Resolution, Time-of-Flight Aerosol Mass Spectrometer,
417 *Anal. Chem.*, 78, 8281-8289, <https://doi.org/10.1021/ac061249n>, 2006.
- 418 Ding, X., Wang, X.-M., Gao, B., Fu, X.-X., He, Q.-F., Zhao, X.-Y., Yu, J.-Z., and Zheng, M.:
419 Tracer-based estimation of secondary organic carbon in the Pearl River Delta, south
420 China, *J. Geophys. Res.-Atmos.*, 117, D05313,
421 <https://doi.org/10.1029/2011JD016596>, 2012.
- 422 Edney, E. O., Driscoll, D. J., Speer, R. E., Weathers, W. S., Kleindienst, T. E., Li, W., and
423 Smith, D. F.: Impact of aerosol liquid water on secondary organic aerosol yields of
424 irradiated toluene/propylene/NOx/(NH₄)₂SO₄/air mixtures, *Atmos. Environ.*, 34,
425 3907-3919, 2000.
- 426 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud
427 droplets and aqueous particles (aqSOA): a review of laboratory, field and model
428 studies, *Atmos. Chem. Phys.*, 11, 11069-11102, [https://doi.org/10.5194/acp-11-](https://doi.org/10.5194/acp-11-11069-2011)
429 11069-2011, 2011.
- 430 Faust, J. A., Wong, J. P. S., Lee, A. K. Y., and Abbatt, J. P. D.: Role of Aerosol Liquid Water
431 in Secondary Organic Aerosol Formation from Volatile Organic Compounds,
432 *Environ. Sci. Technol.*, 51, 1405-1413, <https://doi.org/10.1021/acs.est.6b04700>, 2017.
- 433 Fisseha, R., Dommen, J., Sax, M., Paulsen, D., Kalberer, M., Maurer, R., Höfler, F.,
434 Weingartner, E., and Baltensperger, U.: Identification of Organic Acids in Secondary
435 Organic Aerosol and the Corresponding Gas Phase from Chamber Experiments,
436 *Anal. Chem.*, 76, 6535-6540, <https://doi.org/10.1021/ac048975f>, 2004.
- 437 Guo, H., Xu, L., Bougiatioti, A., Cerully, K. M., Capps, S. L., Hite Jr, J. R., Carlton, A. G.,
438 Lee, S. H., Bergin, M. H., Ng, N. L., Nenes, A., and Weber, R. J.: Fine-particle water
439 and pH in the southeastern United States, *Atmos. Chem. Phys.*, 15, 5211-5228,
440 <https://doi.org/10.5194/acp-15-5211-2015>, 2015.
- 441 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
442 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F.,
443 Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L.,
444 Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
445 Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The
446 formation, properties and impact of secondary organic aerosol: current and emerging
447 issues, *Atmos. Chem. Phys.*, 9, 5155-5236, <https://doi.org/10.5194/acp-9-5155-2009>,
448 2009.
- 449 Hamilton, J. F., Webb, P. J., Lewis, A. C., and Reviejo, M. M.: Quantifying small molecules
450 in secondary organic aerosol formed during the photo-oxidation of toluene with
451 hydroxyl radicals, *Atmos. Environ.*, 39, 7263-7275,
452 <http://dx.doi.org/10.1016/j.atmosenv.2005.09.006>, 2005.
- 453 Heald, C. L., Kroll, J. H., Jimenez, J. L., Docherty, K. S., DeCarlo, P. F., Aiken, A. C., Chen,
454 Q., Martin, S. T., Farmer, D. K., and Artaxo, P.: A simplified description of the
455 evolution of organic aerosol composition in the atmosphere, *Geophys. Res. Lett.*, 37,
456 L08803, <https://doi.org/10.1029/2010gl042737>, 2010.



- 457 Hennigan, C. J., Bergin, M. H., Dibb, J. E., and Weber, R. J.: Enhanced secondary organic
458 aerosol formation due to water uptake by fine particles, *Geophys. Res. Lett.*, 35,
459 L18801, <https://doi.org/10.1029/2008GL035046>, 2008.
- 460 Hildebrandt, L., Donahue, N. M., and Pandis, S. N.: High formation of secondary organic
461 aerosol from the photo-oxidation of toluene, *Atmos. Chem. Phys.*, 9, 2973-2986,
462 <https://doi.org/10.5194/acp-9-2973-2009>, 2009.
- 463 Hodzic, A., Jimenez, J. L., Madronich, S., Canagaratna, M. R., DeCarlo, P. F., Kleinman, L.,
464 and Fast, J.: Modeling organic aerosols in a megacity: potential contribution of semi-
465 volatile and intermediate volatility primary organic compounds to secondary organic
466 aerosol formation, *Atmos. Chem. Phys.*, 10, 5491-5514, <https://doi.org/10.5194/acp-10-5491-2010>, 2010.
- 468 Huang, D. D., Zhang, X., Dalleska, N. F., Lignell, H., Coggon, M. M., Chan, C.-M., Flagan,
469 R. C., Seinfeld, J. H., and Chan, C. K.: A note on the effects of inorganic seed aerosol
470 on the oxidation state of secondary organic aerosol— α -Pinene ozonolysis, *J.*
471 *Geophys. Res.-Atmos.*, 121, 2016JD025999, <https://doi.org/10.1002/2016JD025999>,
472 2016.
- 473 Kamens, R. M., Zhang, H. F., Chen, E. H., Zhou, Y., Parikh, H. M., Wilson, R. L., Galloway,
474 K. E., and Rosen, E. P.: Secondary organic aerosol formation from toluene in an
475 atmospheric hydrocarbon mixture: Water and particle seed effects, *Atmos. Environ.*,
476 45, 2324-2334, <https://doi.org/doi:10.1016/j.atmosenv.2010.11.007>, 2011.
- 477 Kang, E., Root, M. J., Toohey, D. W., and Brune, W. H.: Introducing the concept of Potential
478 Aerosol Mass (PAM), *Atmos. Chem. Phys.*, 7, 5727-5744,
479 <https://doi.org/10.5194/acp-7-5727-2007>, 2007.
- 480 Kang, E., Toohey, D. W., and Brune, W. H.: Dependence of SOA oxidation on organic
481 aerosol mass concentration and OH exposure: experimental PAM chamber studies,
482 *Atmos. Chem. Phys.*, 11, 1837-1852, <https://doi.org/10.5194/acp-11-1837-2011>,
483 2011.
- 484 Krechmer, J. E., Pagonis, D., Ziemann, P. J., and Jimenez, J. L.: Quantification of Gas-Wall
485 Partitioning in Teflon Environmental Chambers Using Rapid Bursts of Low-
486 Volatility Oxidized Species Generated in Situ, *Environ. Sci. Technol.*, 50, 5757-5765,
487 <https://doi.org/10.1021/acs.est.6b00606>, 2016.
- 488 Kreidenweis, S. M., Petters, M. D., and DeMott, P. J.: Single-parameter estimates of
489 aerosol water content, *Environ. Res. Lett.*, 3, 035002, 2008.
- 490 Kroll, J. H., Smith, J. D., Che, D. L., Kessler, S. H., Worsnop, D. R., and Wilson, K. R.:
491 Measurement of fragmentation and functionalization pathways in the heterogeneous
492 oxidation of oxidized organic aerosol, *Phys. Chem. Chem. Phys.*, 11, 8005-8014,
493 <https://doi.org/10.1039/B905289E>, 2009.
- 494 Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K.
495 R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith,
496 J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric for
497 describing the chemistry of atmospheric organic aerosol, *Nat. Chem.*, 3, 133-139,
498 2011.
- 499 Lambe, A. T., Ahern, A. T., Williams, L. R., Slowik, J. G., Wong, J. P. S., Abbatt, J. P. D.,
500 Brune, W. H., Ng, N. L., Wright, J. P., Croasdale, D. R., Worsnop, D. R., Davidovits,



- 501 P., and Onasch, T. B.: Characterization of aerosol photooxidation flow reactors:
502 heterogeneous oxidation, secondary organic aerosol formation and cloud
503 condensation nuclei activity measurements, *Atmos. Meas. Tech.*, 4, 445-461,
504 <https://doi.org/10.5194/amt-4-445-2011>, 2011a.
- 505 Lambe, A. T., Onasch, T. B., Massoli, P., Croasdale, D. R., Wright, J. P., Ahern, A. T.,
506 Williams, L. R., Worsnop, D. R., Brune, W. H., and Davidovits, P.: Laboratory
507 studies of the chemical composition and cloud condensation nuclei (CCN) activity of
508 secondary organic aerosol (SOA) and oxidized primary organic aerosol (OPOA),
509 *Atmos. Chem. Phys.*, 11, 8913-8928, <https://doi.org/10.5194/acp-11-8913-2011>,
510 2011b.
- 511 Lambe, A. T., Chhabra, P. S., Onasch, T. B., Brune, W. H., Hunter, J. F., Kroll, J. H.,
512 Cummings, M. J., Brogan, J. F., Parmar, Y., Worsnop, D. R., Kolb, C. E., and
513 Davidovits, P.: Effect of oxidant concentration, exposure time, and seed particles on
514 secondary organic aerosol chemical composition and yield, *Atmos. Chem. Phys.*, 15,
515 3063-3075, <https://doi.org/10.5194/acp-15-3063-2015>, 2015.
- 516 Lee, A. K. Y., Herckes, P., Leaitch, W. R., Macdonald, A. M., and Abbatt, J. P. D.: Aqueous
517 OH oxidation of ambient organic aerosol and cloud water organics: Formation of
518 highly oxidized products, *Geophys. Res. Lett.*, 38, L11805,
519 <https://doi.org/10.1029/2011GL047439>, 2011.
- 520 Lee, Y. H., and Adams, P. J.: Evaluation of aerosol distributions in the GISS-TOMAS global
521 aerosol microphysics model with remote sensing observations, *Atmos. Chem. Phys.*,
522 10, 2129-2144, <https://doi.org/10.5194/acp-10-2129-2010>, 2010.
- 523 Liao, H., and Seinfeld, J. H.: Global impacts of gas-phase chemistry-aerosol interactions on
524 direct radiative forcing by anthropogenic aerosols and ozone, *J. Geophys. Res.-*
525 *Atmos.*, 110, D18208, <https://doi.org/10.1029/2005JD005907>, 2005.
- 526 Lim, Y. B., Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Aqueous chemistry and
527 its role in secondary organic aerosol (SOA) formation, *Atmos. Chem. Phys.*, 10,
528 10521-10539, <https://doi.org/10.5194/acp-10-10521-2010>, 2010.
- 529 Liu, T., Li, Z., Chan, M., and Chan, C. K.: Formation of secondary organic aerosols from gas-
530 phase emissions of heated cooking oils, *Atmos. Chem. Phys.*, 17, 7333-7344,
531 <https://doi.org/10.5194/acp-17-7333-2017>, 2017.
- 532 Mahmud, A., and Barsanti, K.: Improving the representation of secondary organic aerosol
533 (SOA) in the MOZART-4 global chemical transport model, *Geosci. Model Dev.*, 6,
534 961-980, [10.5194/gmd-6-961-2013](https://doi.org/10.5194/gmd-6-961-2013), 2013.
- 535 Mao, J., Ren, X., Brune, W. H., Olson, J. R., Crawford, J. H., Fried, A., Huey, L. G., Cohen,
536 R. C., Heikes, B., Singh, H. B., Blake, D. R., Sachse, G. W., Diskin, G. S., Hall, S.
537 R., and Shetter, R. E.: Airborne measurement of OH reactivity during INTEX-B,
538 *Atmos. Chem. Phys.*, 9, 163-173, <https://doi.org/10.5194/acp-9-163-2009>, 2009.
- 539 Matsunaga, A., and Ziemann ‡, P. J.: Gas-Wall Partitioning of Organic Compounds in a
540 Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield
541 Measurements, *Aerosol Sci. Tech.*, 44, 881-892,
542 <https://doi.org/10.1080/02786826.2010.501044>, 2010.
- 543 Matthew, B. M., Middlebrook, A. M., and Onasch, T. B.: Collection Efficiencies in an
544 Aerodyne Aerosol Mass Spectrometer as a Function of Particle Phase for Laboratory



- 545 Generated Aerosols, *Aerosol Sci. Tech.*, 42, 884-898,
546 <https://doi.org/10.1080/02786820802356797>, 2008.
- 547 Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of
548 Composition-Dependent Collection Efficiencies for the Aerodyne Aerosol Mass
549 Spectrometer using Field Data, *Aerosol Sci. Tech.*, 46, 258-271,
550 <https://doi.org/10.1080/02786826.2011.620041>, 2012.
- 551 Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.:
552 Secondary organic aerosol formation from m-xylene, toluene, and benzene, *Atmos.*
553 *Chem. Phys.*, 7, 3909-3922, <https://doi.org/10.5194/acp-7-3909-2007>, 2007.
- 554 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H.,
555 Docherty, K. S., Chhabra, P. S., Bahreini, R., Murphy, S. M., Seinfeld, J. H.,
556 Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H., Dinar,
557 E., Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in
558 Northern Hemispheric datasets from Aerosol Mass Spectrometry, *Atmos. Chem.*
559 *Phys.*, 10, 4625-4641, <https://doi.org/10.5194/acp-10-4625-2010>, 2010.
- 560 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H., and Worsnop,
561 D. R.: Changes in organic aerosol composition with aging inferred from aerosol mass
562 spectra, *Atmos. Chem. Phys.*, 11, 6465-6474, [https://doi.org/10.5194/acp-11-6465-](https://doi.org/10.5194/acp-11-6465-2011)
563 2011, 2011.
- 564 Nguyen, T. K. V., Zhang, Q., Jimenez, J. L., Pike, M., and Carlton, A. G.: Liquid Water:
565 Ubiquitous Contributor to Aerosol Mass, *Environ. Sci. Technol. Lett.*, 3, 257-263,
566 <https://doi.org/10.1021/acs.estlett.6b00167>, 2016.
- 567 Parikh, H. M., Carlton, A. G., Vizuete, W., and Kamens, R. M.: Modeling secondary organic
568 aerosol using a dynamic partitioning approach incorporating particle aqueous-phase
569 chemistry, *Atmos. Environ.*, 45, 1126-1137,
570 <http://dx.doi.org/10.1016/j.atmosenv.2010.11.027>, 2011.
- 571 Peng, Z., Day, D. A., Ortega, A. M., Palm, B. B., Hu, W., Stark, H., Li, R., Tsigaridis, K.,
572 Brune, W. H., and Jimenez, J. L.: Non-OH chemistry in oxidation flow reactors for
573 the study of atmospheric chemistry systematically examined by modeling, *Atmos.*
574 *Chem. Phys.*, 16, 4283-4305, <https://doi.org/10.5194/acp-16-4283-2016>, 2016.
- 575 Romonosky, D. E., Laskin, A., Laskin, J., and Nizkorodov, S. A.: High-Resolution Mass
576 Spectrometry and Molecular Characterization of Aqueous Photochemistry Products
577 of Common Types of Secondary Organic Aerosols, *The Journal of Physical*
578 *Chemistry A*, 119, 2594-2606, <https://doi.org/10.1021/jp509476r>, 2015.
- 579 Rudich, Y., Donahue, N. M., and Mentel, T. F.: Aging of organic aerosol: Bridging the gap
580 between laboratory and field studies, *Annual Review of Physical Chemistry*, 58, 321-
581 352, <https://doi.org/10.1146/annurev.physchem.58.032806.104432>, 2007.
- 582 Ruehl, C. R., Nah, T., Isaacman, G., Worton, D. R., Chan, A. W. H., Kolesar, K. R., Cappa,
583 C. D., Goldstein, A. H., and Wilson, K. R.: The Influence of Molecular Structure and
584 Aerosol Phase on the Heterogeneous Oxidation of Normal and Branched Alkanes by
585 OH, *J. Phys. Chem. A*, 117, 3990-4000, <https://doi.org/10.1021/jp401888q>, 2013.
- 586 Sareen, N., Waxman, E. M., Turpin, B. J., Volkamer, R., and Carlton, A. G.: Potential of
587 Aerosol Liquid Water to Facilitate Organic Aerosol Formation: Assessing Knowledge



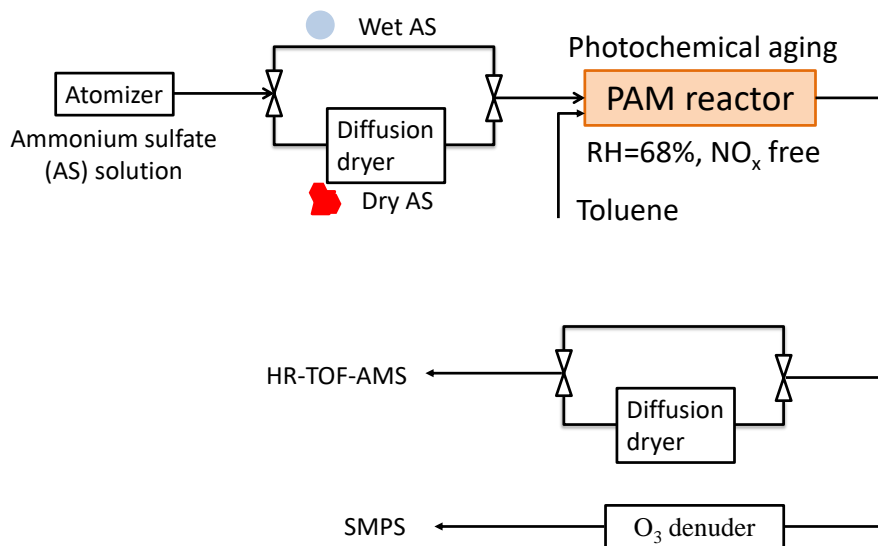
- 588 Gaps about Precursors and Partitioning, *Environ. Sci. Technol.*, 51, 3327-3335,
589 <https://doi.org/10.1021/acs.est.6b04540>, 2017.
- 590 Sato, K., Hatakeyama, S., and Imamura, T.: Secondary organic aerosol formation during the
591 photooxidation of toluene: NO_x dependence of chemical composition, *J. Phys. Chem.*
592 *A*, 111, 9796-9808, <https://doi.org/10.1021/jp071419f>, 2007.
- 593 Seinfeld, J. H.; Pandis, S. N. *Atmospheric Chemistry and Physics: From Air Pollution to*
594 *Climate Change*, 2nd ed.; Wiley: Hoboken, NJ, 2006.
- 595 Smith, M. L., Kuwata, M., and Martin, S. T.: Secondary Organic Material Produced by the
596 Dark Ozonolysis of α -Pinene Minimally Affects the Deliquescence and Efflorescence
597 of Ammonium Sulfate, *Aerosol Sci. Tech.*, 45, 244-261,
598 <https://doi.org/10.1080/02786826.2010.532178>, 2011.
- 599 Smith, M. L., Bertram, A. K., and Martin, S. T.: Deliquescence, efflorescence, and phase
600 miscibility of mixed particles of ammonium sulfate and isoprene-derived secondary
601 organic material, *Atmos. Chem. Phys.*, 12, 9613-9628, [https://doi.org/10.5194/acp-](https://doi.org/10.5194/acp-12-9613-2012)
602 [12-9613-2012](https://doi.org/10.5194/acp-12-9613-2012), 2012.
- 603 Smith, M. L., You, Y., Kuwata, M., Bertram, A. K., and Martin, S. T.: Phase Transitions and
604 Phase Miscibility of Mixed Particles of Ammonium Sulfate, Toluene-Derived
605 Secondary Organic Material, and Water, *J. Phys. Chem. A*, 117, 8895-8906,
606 <https://doi.org/10.1021/jp405095e>, 2013.
- 607 Song, M., Liu, P. F., Hanna, S. J., Zaveri, R. A., Potter, K., You, Y., Martin, S. T., and
608 Bertram, A. K.: Relative humidity-dependent viscosity of secondary organic material
609 from toluene photo-oxidation and possible implications for organic particulate matter
610 over megacities, *Atmos. Chem. Phys.*, 16, 8817-8830, [https://doi.org/10.5194/acp-16-](https://doi.org/10.5194/acp-16-8817-2016)
611 [8817-2016](https://doi.org/10.5194/acp-16-8817-2016), 2016.
- 612 Takahama, S., Pathak, R. K., and Pandis, S. N.: Efflorescence Transitions of Ammonium
613 Sulfate Particles Coated with Secondary Organic Aerosol, *Environ. Sci. Technol.*, 41,
614 2289-2295, <https://doi.org/10.1021/es0619915>, 2007.
- 615 Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D., Molina, L.
616 T., Worsnop, D. R., and Molina, M. J.: Secondary organic aerosol formation from
617 anthropogenic air pollution: Rapid and higher than expected, *Geophys. Res. Lett.*, 33,
618 L17811, <https://doi.org/10.1029/2006gl026899>, 2006.
- 619 Wong, J. P. S., Lee, A. K. Y., and Abbatt, J. P. D.: Impacts of Sulfate Seed Acidity and Water
620 Content on Isoprene Secondary Organic Aerosol Formation, *Environ. Sci. Technol.*,
621 49, 13215-13221, <https://doi.org/10.1021/acs.est.5b02686>, 2015.
- 622 Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
623 Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of
624 secondary organic aerosol, *P. Natl. Acad. Sci.*, 111, 5802-5807,
625 <https://doi.org/10.1073/pnas.1404727111>, 2014.
- 626 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C., and
627 Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, *Atmos. Chem. Phys.*, 15,
628 4197-4214, <https://doi.org/10.5194/acp-15-4197-2015>, 2015.
- 629 Zhang, Z., Zhang, Y., Wang, X., Lü, S., Huang, Z., Huang, X., Yang, W., Wang, Y., and
630 Zhang, Q.: Spatiotemporal patterns and source implications of aromatic hydrocarbons



631 at six rural sites across China's developed coastal regions, *J. Geophys. Res.-Atmos.*,
632 121, 2016JD025115, <https://doi.org/10.1002/2016JD025115>, 2016.
633 Zhao, Y., Saleh, R., Saliba, G., Presto, A. A., Gordon, T. D., Drozd, G. T., Goldstein, A. H.,
634 Donahue, N. M., and Robinson, A. L.: Reducing secondary organic aerosol formation
635 from gasoline vehicle exhaust, *P. Natl. Acad. Sci.*, 114, 6984-6989,
636 <https://doi.org/10.1073/pnas.1620911114>, 2017.
637

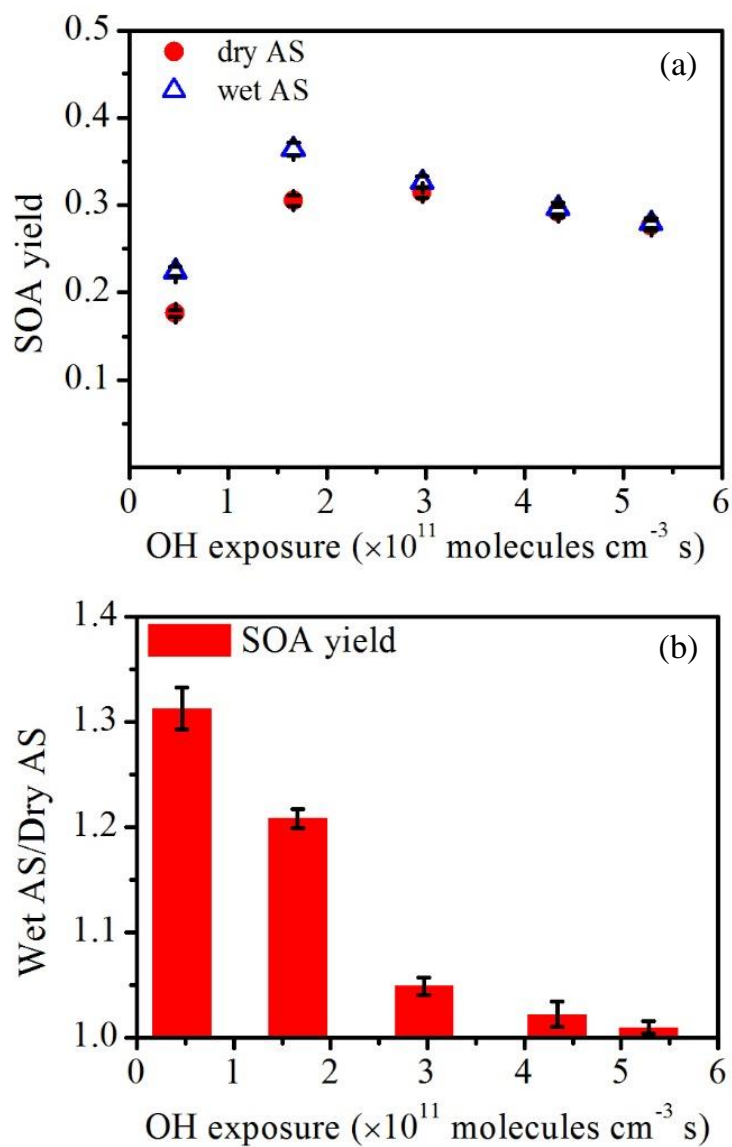


638



639

640 **Fig. 1.** Schematic of the experimental setup. The aqueous ammonium sulfate (AS) seed
641 particles either passed through a diffusion dryer so that the phase of the seed particles
642 could be altered or bypassed the diffusion dryer. Either wet or dry AS served as seed
643 particles for the experiments.



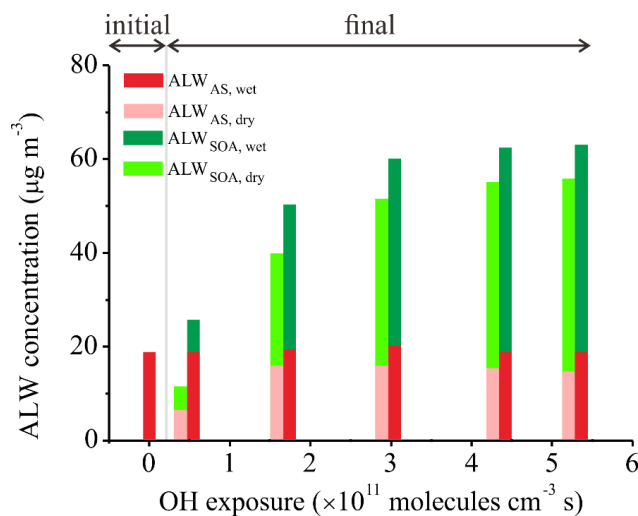
644

645 **Fig. 2.** (a) Yield of toluene-derived SOA formed on initially wet and dry AS as a

646 function of OH exposure. (b) Ratio of SOA yields on initially wet AS to those on

647 initially dry AS as a function of OH exposure.

648



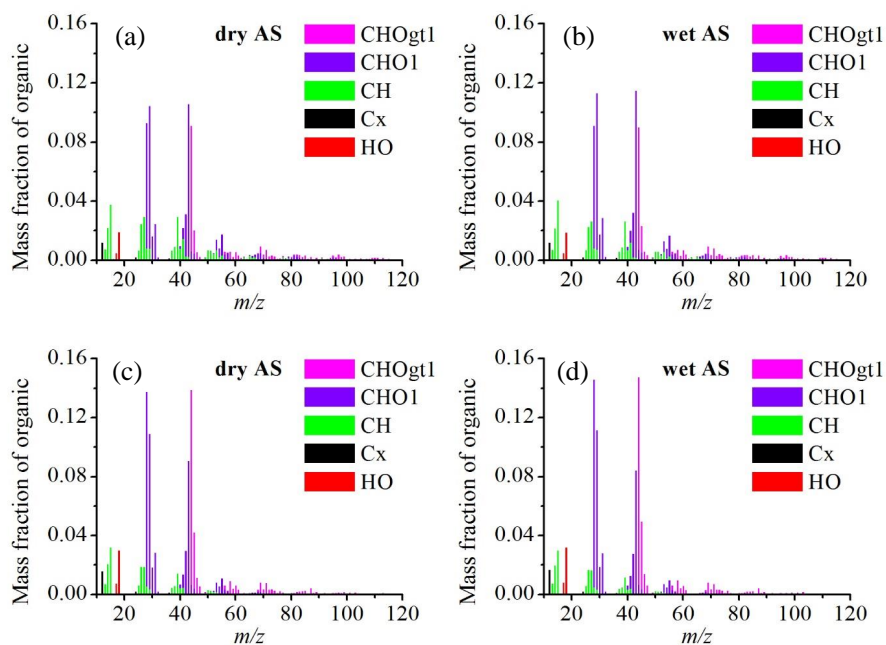
649

650 **Fig. 3.** Mass concentration of ALW uptake by AS and toluene-derived SOA before

651 (initial) and after reactions (final) for both initially wet and dry AS seeds. Adjoining

652 bars for initially wet and dry seeds have same OH exposures.

653

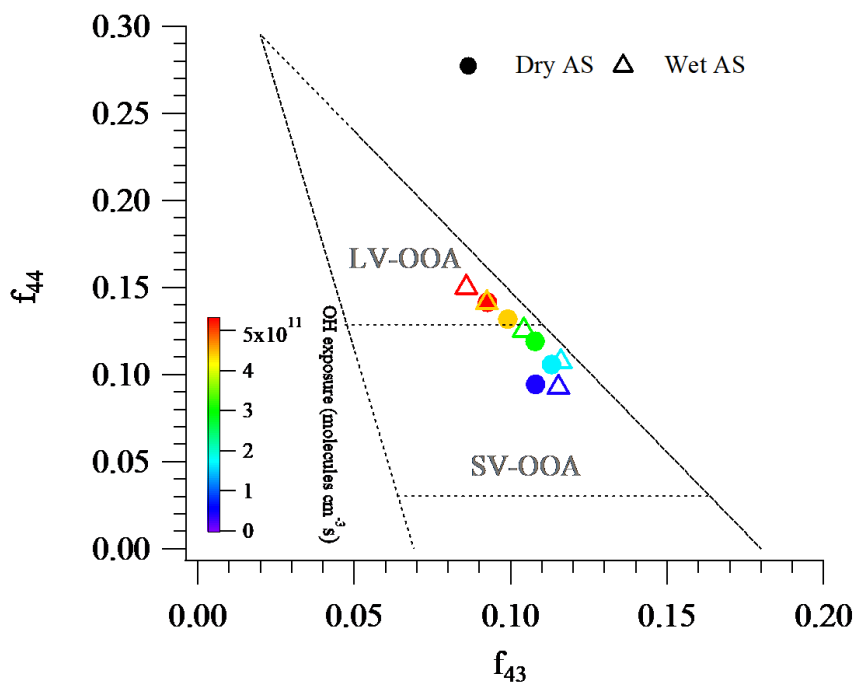


654

655 **Fig. 4.** High-resolution mass spectra of toluene-derived SOA on initially wet and dry

656 AS at an OH exposure of (a, b) 4.66×10^{10} molecules cm^{-3} s and (c, d) 5.28×10^{11}

657 molecules cm^{-3} s.



658

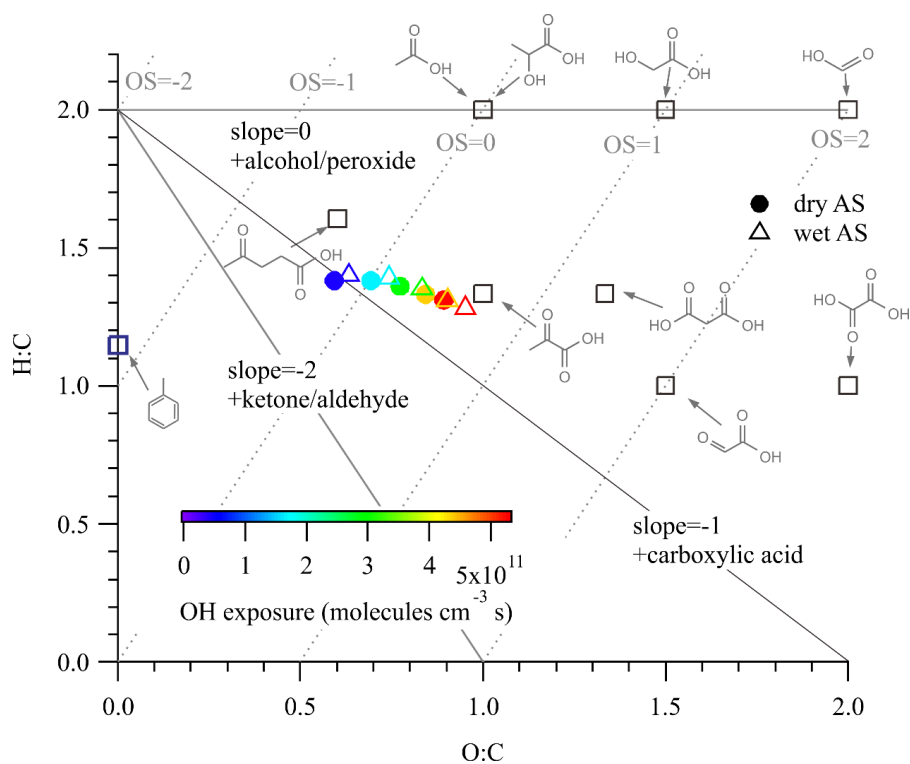
659 **Fig. 5.** Fractions of total organic signal at m/z 43 (f_{43}) vs. m/z 44 (f_{44}) from SOA data

660 obtained in this study together with the triangle plot of Ng et al. (2010). Ambient SV-

661 OOA and LV-OOA regions are adapted from Ng et al. (2010). Data are colored

662 according to the OH exposure.

663



664

665 **Fig. 6.** Van Krevelen diagram of SOA derived from the photooxidation of toluene on
 666 initially wet and dry AS seed particles. SOA data are colored according to the OH
 667 exposure. Products identified in toluene-derived SOA are shown in boxes (Bloss et al.,
 668 2005; Hamilton et al., 2005; Sato et al., 2007). Average carbon oxidation states from
 669 Kroll et al. (2011) and functionalization slopes from Heald et al. (2010) are shown for
 670 reference.

671