Hygroscopic Properties of Aminium Sulphate Aerosols

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Abstract. Alkylaminium sulphates originate from the neutralisation reaction between short-chained amines and sulphuric acid and have been detected in atmospheric aerosol particles. Their physicochemical behaviour is less well characterised than their

- 15 inorganic equivalent, ammonium sulphate, even though they play a role in atmospheric processes such as the nucleation and growth of new particles and cloud droplet formation. In this work, a comparative evaporation kinetics experimental technique using a cylindrical electrodynamic balance is applied to determine the hygroscopic properties of six short-chained alkylaminium sulphates, specifically mono-, di- and tri-methylaminium sulphate and mono-, di- and tri-ethyl aminium sulphate. This approach allows the retrieval of a water activity dependent growth curve in less than 10 s, avoiding the
- 20 uncertainties that can arise from the volatilisation of semi-volatile components. Measurements are made on particles $>5 \mu m$ in radius, avoiding the need to correct equilibrium measurements for droplet surface curvature with assumed values of the droplet surface tension. Variations in equilibrium solution droplet composition with varying water activity are reported over the range 0.5 to >0.98, along with accurate parameterisations of solution density and refractive index. The uncertainties in water activities associated with the hygroscopicity measurements are typically <±0.2% at water activities >0.9 and ∼±1% below 0.9, with
- 25 maximum uncertainties in diameter growth factors of $\pm 0.7\%$. Comparison with previously reported measurements show deviation across the whole water activity range.

1. Introduction

Quantifying the response of aerosol particles to variations in relative humidity (RH) in the atmosphere is important for predicting the impact of aerosols on climate through both direct and indirect radiative forcings (Albrecht, 1989; Kanakidou et al., 2005; Kolb et al., 2010), for understanding the mechanisms of heterogeneous chemistry (Dennis-Smither et al., 2014; Lee

- 5 et al., 2012) and the partitioning of products in the condensed phase (Dusek et al., 2006; Topping et al., 2013a), and for inferring the deposition and loss mechanism of aerosols on surfaces and on inhalation into the respiratory tract (Broday and Georgopoulos, 2001; Haddrell et al., 2015). With an increase in RH, solution droplets must absorb water to maintain an equilibrium balance of water between the gas and condensed phases, leading to the dilution of solutes, increase in mass of condensed phase water and hygroscopic growth in particle size. The hygroscopic response of a wide range of inorganic and
- 10 organic solutes of varying solubility often present as complex mixtures, must be quantified. Thus, the influence of interactions between multiple solutes of varying charge, acidity and hydrophilicity must be understood through predictions of activity coefficients that reflect the departure of solution composition from ideality (Dutcher et al., 2013; Shiraiwa et al., 2013).

Measurements of the hygroscopic growth of ambient aerosol are routinely made: the extent of growth at high RH or the critical

- 15 supersaturation for activation as cloud condensation nuclei are often represented by a single parameter, such as the parameter in κ-Köhler theory (Petters and Kreidenweis, 2007). Resorting to such a reduced parameter model is a consequence of the complex chemical composition of ambient aerosol and the intractability of providing accurate predictions based on a detailed chemical speciation. However, laboratory measurements are essential for exploring the complex details often obscured in measurements of ambient aerosol. These include the roles of pH, liquid-liquid phase separation and mixing state (Topping et
- 20 al., 2013c; You et al., 2014) and the co-condensation of semi-volatile organic species with varying RH (Topping et al., 2013b). When considered alongside measurements of hygroscopic response and change in particle composition with RH for aerosol of well-known chemical composition, detailed predictive tools can be benchmarked and improved, providing a crucial framework for interpreting field measurements. We present here measurements of the hygroscopic response of a class of compounds, aminium sulphates, with the objective of providing accurate data to refine predictive tools of their equilibrium hygroscopic
- 25 response.

Low molecular weight amines are mostly emitted into the atmosphere as gaseous compounds by a number of different sources, both anthropogenic (animal husbandry, food processing and cooking, combustion, pesticides) and natural (oceans, biomass burning, vegetation) (Ge et al., 2011a). These chemical species are basic (with pK_a values from 9.8 to 10.84) (Lide, 2005),

30 highly water-soluble and have high pure liquid vapour pressures (up to hundreds of kPa at 298.15 K) (Ge et al., 2011b). The ambient concentrations of amines in the gas phase can span wide ranges, depending on the sampling location. For example, concentrations can be up to 140 mg m⁻³ close to a city market (Namieśnik et al., 2003), 110-300 ng m⁻³ in the exhaust gas of a waste disposal site (Kallinger and Niessner, 1999), and of the order of tens of µg m⁻³ inside livestock buildings (Kallinger and Niessner, 1999). Concentrations also depend on the season: for example, single amines in the gas phase at a rural site in Turkey have been reported to be in the range 0.92 -7.4 ng m³ in Winter and 0.29-5.16 ng m³ in Summer(Akyüz, 2008);. Further, Pratt et al. (2009) measured a seasonal volatility dependence of alkylamines depending on the particles pH that affects the amines concentrations in the gas phase. In the condensed phase, amines can account for hundreds of pg $m⁻³$ or a few ng $m⁻³$ of

5 aerosol mass. An extensive review of the measured concentrations of a large number of amines both in the gas phase and in aerosols at a large number of rural and urban environments is given by Ge et al. (2011a, 2011b).

The role of amines in new particles formation and growth has been highlighted by computational studies (Loukonen et al., 2010; DePalma et al., 2012; Ortega et al., 2012), as well as by laboratory (Wang et al., 2010b; Almeida et al., 2013) and field

- 10 measurements (Mäkelä et al., 2001; Smith et al., 2010; Kulmala et al., 2013). As an example, trimethylamine was found to enhance the formation and growth of new particles (Wang et al., 2010a) because of the favourable heterogeneous neutralisation reactions between the amine gaseous molecules and $H_2SO_4-H_2O$ clusters. Smith et al. (2010) found variable but considerable concentrations of protonated amines in nanoparticles (8-10 nm diameter) during new particles formation events (47% of detected positive ions at an urban site in Mexico, 23% at remote site in Finland and 10-35% at sampling sites in Atlanta and
- 15 Boulder). Since newly formed secondary particles are estimated to contribute to 45% of cloud condensation nuclei (CCN) (Merikanto et al., 2009), the presence of amines in CCN and their hygroscopic properties need to be taken into account to improve our understanding of the indirect effects of aerosol particles on climate (McFiggans et al., 2005), a key motivator for providing refined characterisation of the hygroscopic growth of aminium salt particles in this work.
- 20 In addition, despite the volatility of short-chained alkylamines, these compounds can undergo gas-to-particle partitioning due to a variety of chemical processes (Ge et al., 2011a): direct solubilisation, oxidation reactions that lead to the formation of secondary organic aerosols, acid-base reactions similar to those of ammonia (Seinfeld and Pandis, 2006) with both inorganic (sulphuric, nitric and hydrochloric acids) and organic acids (Lavi et al., 2015; Liu et al., 2012), and displacement reactions of ammonium cations with aminium cations (Bzdek et al., 2010; Qiu and Zhang, 2013). The product of these last two types of
- 25 chemical reactions and the salts formed by protonated amines following to new particles formation are aminium sulphates, which result from the neutralisation of sulphuric acid and short-chained alkylamines:

$$
H_2SO_4 + 2R_3N \rightleftarrows (R_3NH)_2SO_4 \tag{1}
$$

where R can be either $-CH_3$, $-CH_2CH_3$ or $-H$. The formation of aminium sulphates and other similar aminium salts increases the solubility of short-chained alkylamines from 7-45 wt% for the unprotonated form (calculation from data in Ge et al.

30 (2011b)) to 84-91 wt% for five aminium sulphates in Clegg et al. (2013). This, correspondingly, increases their partitioning from the gas to the condensed phase (Barsanti et al., 2009; Yli-Juuti et al., 2013). However, the physicochemical properties of aminium sulphates are much less well characterised than their inorganic counterpart, $(NH₄)₂SO₄$, even though they can play a fundamental role in the nucleation and growth of new particles (DePalma et al., 2012; Loukonen et al., 2010; Wang et al., 2010b) and in cloud formation (Lavi et al., 2013). Some recent work has attempted to fill this gap, reporting the densities (Clegg et al., 2013; Lavi et al., 2013; Qiu and Zhang, 2012), CCN activity, optical properties (Lavi et al., 2013), hygroscopicity and phase transitions of aminium sulphates (Chu et al., 2015; Clegg et al., 2013; Lavi et al., 2013; Qiu and Zhang, 2012; Sauerwein et al., 2015), specifically for the methyl and ethylaminium sulphates (or subsets of these six compounds). In this work, we report measurements of the hygroscopic behaviour of the three methylaminium sulphates (methylaminium sulphate,

- 5 MMAS; dimethylaminium sulphate, DMAS; trimethylaminium sulphate, TMAS) and the three ethylaminium sulphates (ethylaminium sulphate, MEAS; diethylaminium sulphate, DEAS; triethylaminium sulphate, TEAS) from water evaporation experiments made using a comparative kinetic Electrodynamic Balance (CK-EDB) (Rovelli et al., 2016). In Section 2 we review the experimental details and procedures for performing hygroscopic growth measurements using the CK-EDB. In Section 3, we present measurements of the hygroscopic growth of the six aminium sulphate salts, comparing our measurements
- 10 with literature values and considering the accuracy and reproducibility of measurements with the CK-EDB.

2 Experimental

We first describe the CK-EDB technique and the data analysis before describing the method for preparing aqueous starting solutions of the aminium sulphates.

2.1 Hygroscopic Properties from Comparative Kinetics Measurements in a CK-EDB

- 15 Electrodynamic trapping of single charged droplets has been used to investigate various properties of confined particles, including optical properties (Barnes et al., 1997), vapour pressures of low volatility compounds (Pope et al., 2010), hygroscopic properties of atmospherically-relevant aqueous solutions (Choi and Chan, 2002; Chu et al., 2015; Peng et al., 2001; Rickards et al., 2013) and of pharmaceutical aerosols (Haddrell et al., 2013; Peng et al., 2000), and evaporation dynamics of aqueous droplets (Heinisch et al., 2009; Shulman et al., 1997; Zobrist et al., 2011). The experimental setup used in this work has been
- 20 presented in previous publications (Davies, 2014; Davies et al., 2012a, 2012b, 2013; Haddrell et al., 2012; Miles et al., 2012) and a schematic of it can be found in the Supplementary Information (Fig. S1). In particular, the approach used to quantify the hygroscopic properties of single confined particles from comparative kinetics measurements using a CK-EDB has also been previously discussed (Davies et al., 2013) and its application over a wide solution water activity range from 0.5 to >0.99 has been extensively validated (Rovelli et al., 2016).

In a CK-EDB single charged droplets from two solutions with known chemical composition and concentration can be sequentially dispensed on-demand by means of two alternatively operated microdispensers and trapped within the electrodynamic field generated by a set of concentric cylindrical electrodes. This electrode geometry guarantees a stable and tight trapping of single droplets within 100 ms from their generation. Single droplets are confined in the trapping chamber 30 within a nitrogen stream with controlled mass flow, temperature and relative humidity. The gas flow RH is modified by mixing

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different ratios of a humidified and a dry nitrogen flow and is inferred from the evaporation kinetics of probe droplets, as described below. The temperature within the trapping chamber is controlled by a circulating a 1:1 volume mixture of water and ethylene glycol, through the lid and the bottom of the chamber. Temperature and gas phase RH ranges that are accessible with this experimental setup are -25 to 50 $^{\circ}$ C and 0 to 99%, respectively. All the comparative evaporation kinetics

5 measurements presented here were performed at 20 °C and at gas phase RH values between ~50-90%. Trapped droplets are illuminated by laser light (532 nm) and the resulting elastic scattering light pattern is collected every 0.01 s by means of a CCD camera and used to keep track of changes in size of a single levitated droplet using the geometrical optics approximation (Glantschnig and Chen, 1981). Note that the variations in refractive index due to water evaporation from each droplet are taken into account for an accurate determination of droplet size, as described in a previous publication (Davies et al., 2012b).

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The evaporation rate of water from a droplet containing one or more solutes at a certain RH and temperature depends on the hygroscopic properties of the solution. In comparative kinetics experiments, the evaporation rate of single probe droplets with known composition and known evaporation kinetics (either pure water or a NaCl solution) is compared to that of single sample droplets in order to determine the hygroscopic properties of the solution droplet being studied (Davies et al., 2013; Rovelli et

15 al., 2016). To do so, typical CK-EDB experiment are composed of a series of at least ten alternating single probe and sample droplets (Fig. 1a).

The retrieval of the hygroscopic properties from the radius (*a*) vs. time evaporation profiles of a sequence of singly-trapped probe and sample droplets is shown schematically in Fig. 1. Different analyses are carried out for probe (Fig. 1b) and sample

- 20 droplets (Fig. 1c). Either pure water (as shown in Fig. 1a) or NaCl solution droplets with known initial salt concentration can be used as a probe. The evaporation profiles of the probe droplets are compared with simulations obtained from the evaporation/condensation kinetics model by Kulmala et al. (1993). By doing so, the gas phase RH can be inferred, either from fitting the a^2 vs. time evaporation profile of pure water droplets or from the equilibrated radius of the NaCl solution droplets. Both methods, together with their associated uncertainties and their RH ranges of applicability, have been discussed previously
- 25 (Davies et al., 2013; Rovelli et al., 2016). RH fluctuations over the run of ten pairs or more of probe and sample droplets are very slight, typically of the order of 0.2% RH and never exceeding 0.5% RH. However, it should be noted that slight RH fluctuations are taken into account in our approach: the gas phase RH is monitored before every sample droplet by injecting a probe droplet and data from this probe droplet are directly used in the sample droplet evaporation analysis..
- 30 For the analysis of the radius (*a*) vs. time data of sample droplets, the radius data (Fig. 1c, inset panel i) are converted to droplet mass (inset ii) by using a $3rd$ order polynomial parameterisation of density as a function of the solute mass fraction, which is discussed below. The mass flux of water leaving the droplet during evaporation (d*m*/d*t*) is then calculated and, knowing the gas phase RH from the probe droplet analysis, Kulmala's equations for evaporation kinetics are applied to calculate the temporal variation of water activity (a_w) in the droplet (inset iii). In addition, knowing the initial concentration and size of the

droplet at generation, the radius of the dry particle (a_{dry}) can be estimated and a radial growth factor $(GF_r = a / a_{dry})$ can be calculated for each of the measured radii (inset iv). Results from inset panels (iii) and (iv) are then combined and typical *GF^r* vs *a^w* growth curves are obtained. The key thermodynamic quantities that describe the hygroscopic properties of the tested solution (moles of water per mole of solute in solution, $n_{\text{water}}/n_{\text{solute}}$; osmotic coefficients, ϕ_{st}) can be calculated if the densities

- 5 of the aqueous solutions are also known (see below). Details of the treatment of the experimental uncertainties and their influence on each of the computed quantities are described in the Supplementary Information (Table S1). When error bars are not shown in the figures in the following sections, the reader can assume that they are smaller than the size of the corresponding data point.
- 10 A previous study (Cai et al., 2016) showed that the molar refraction mixing rule, together with a $3rd$ order polynomial parameterisation of density as a function of the square-rooted mass fraction of solute (*mfs*), represent the best approach to predicting refractive indices (*m*) and densities (*ρ*) of solutions of organic compounds for which bulk data of such quantities is available for solute mass fractions up to at least 0.4. In this work the densities of at least ten solutions with different concentrations for each aminium sulphate were measured with a density meter (Densito 30PX, Mettler Toledo, accuracy of
- ± 0.001 g cm⁻³, calibrated with pure water before each use). Densities were measured at ambient temperature, which varies in the laboratory between 293 K and 295 K; temperatures were always registered together with the measured density values. In addition, refractive indices of the same solutions were measured at 589 nm by means of a refractometer (Palm Abbe II, Misco, precision of ± 0.0001 , calibration with pure water before each use). The measured density and *m* values for each aminium sulphate solution are provided in the Supplementary Information (Table S2), together with the $3rd$ order polynomial and the
- 20 molar refraction mixing rule fittings for each compound (Table S3 and Fig. S2). A brief description of the molar refraction mixing rule application is also provided in the Supplementary Information. These data have been presented previously and discussed by Cai et al. (2016) along with measurements from a large number of organic aqueous solutions.

2.2 Preparation of the Solutions

- 25 Aminium sulphate stock solutions were prepared by the neutralization of solutions of each of the six amines with aqueous sulphuric acid. The commercial amines stock solutions (Sigma Aldrich, MMA, \sim 40 wt%; DMA, \sim 40 wt%; TMA, \sim 45 wt%; MEA, \sim 66.0-72.0 wt%; DEA \geq 99.5 wt%; TEA \geq 99 wt%) were titrated with standardised HCl (1 M, SLS) to determine their mass concentrations accurately. Three repetitions were performed for each titration and the pH was measured throughout by means of a pH-meter (HI 8314, Hanna Instruments), which was calibrated with standard pH 7 and pH 4 solutions. Before
- 30 the titration, the amine stock solutions were always diluted down to 1-5 wt% in order to minimize the heat generated by the neutralisation reaction and to minimise volatilisation of the amine. In addition, the HCl (and later the H_2SO_4) used for the standardisation of the commercial amine solutions were titrated with $Na_2CO_3 \ (geq 99.5\%$, Alfa Aesar), which was first dried at 225°C for 3 h before weighing to make sure that no water was adsorbed on it. Three repetitions were performed in these cases.

For the preparation of the aminium sulphates stock solutions, stoichiometric amounts of the standardised H_2SO_4 and amine solutions were mixed with an initial concentration of both solutions around 40 wt%. The pH of the solution mixture was monitored for the whole duration of the reaction to ensure that all of the amine in the solution had reacted. The concentrated

- 5 stock solutions of the salts that result from this procedure were subsequently diluted down to a weight percentage of \sim 5%, in order to obtain a suitable starting concentration for the CK-EDB comparative kinetics measurements. During both the titration of the amine stock solution with HCl and the preparation of the aminium sulphates solutions with H_2SO_4 , the amine solution was kept in an ice bath (0 °C) and the addition of the acid was performed slowly and dropwise, in order to dissipate the heat generated by the neutralisation reaction and to avoid any possible amine volatilization. Phase separation was observed when
- 10 titrating the TEA commercial solution and its concentration was determined to be 81.6 wt% (0.52 wt% standard deviation over 3 repetitions), which is considerably lower than the ≥ 99 wt% concentration value given by the manufacturer. This is possibly due to the much lower solubility of TEA $(0.7 \text{ mol kg}^{-1})$ than the other amines (Ge et al., 2011b), leading to incomplete solvation of the amine in water and an inaccurate measurement of pH during the titration. Thus, we instead assumed that the commercial TEA solution was a 99 wt% concentration and note that the results for the TEA system should be interpreted with some caution.
- 15 This experimental procedure ensured that the concentrations of the reagents were well known and, consequently, that the concentrations of the stock solutions prepared for aminium sulphates were similarly well known. The uncertainties in the commercial solution concentrations of the reagents, as determined from the repeated titrations, were taken into account for the calculation of the overall experimental uncertainties indicated in the Supplementary Information (Table S1).
- 20 To validate the procedure for solution preparation described above, the steps were carried out for the preparation of $(NH₄)₂SO₄$ from the reaction of ammonia and sulphuric acid. The resulting salt solution was then used in comparative kinetics measurements and the hygroscopicity of the ammonium sulphate from reaction was compared with calculations from the Extended Aerosol Inorganics Model (E-AIM) (Wexler and Clegg, 2002). Figure 2 shows the hygroscopic properties of (NH4)2SO⁴ droplets prepared in this way from the direct reaction of ammonia and sulphuric acid, reporting values of $n_{\text{water}}/n_{\text{solute}}$, vs. a_w (Panel (a)) and osmotic coefficients (ϕ_{st}) plotted against the square-root of the sulphate molality (*m*(SO₄²) 0.5 , Panel (b)). Osmotic coefficients are useful parameters to represent the deviation of a solution from an ideal behaviour and they are defined as indicated in Eq. (2):

$$
\phi_{st} = -\frac{\ln(a_w)}{M_w 3m/1000} \tag{2}
$$

where M_w is the molecular weight of water, *m* is the molality of the solute and 3 is the stoichiometric number of ions in the 30 salt. The value of *ϕst* tends to 1.0 in the limit of an infinitely dilute solution, in accordance with the Debye-Hückel limiting law (Robinson and Stokes, 1970).

In a previous publication (Rovelli et al., 2016) we showed that it is possible to achieve very good agreement with predictions from the E-AIM model for well-characterised inorganic compounds with the CK-EDB experimental technique within an uncertainty in a_w of ± 0.002 . The plots in Fig. 2 show the averaged data obtained from two datasets of ten droplets of aqueous (NH_4) ₂SO₄ (black and open circles). In Fig. 2b the effect of a ± 0.002 error on a_w on the modelled osmotic coefficients values

- 5 is shown with dashed lines. The osmotic coefficients agree well with calculations from the E-AIM model and lie within the envelope associated with this previous estimate of typical experimental uncertainty for the CK-EDB technique. This demonstrates that the volatility of ammonia is not a significant problem when the neutralisation reaction with H2SO⁴ is performed. In addition, since the vapour pressure of ammonia (1956 kPa at 298.15 K) (Lide, 2005) is even higher than the vapour pressure of the most volatile of the six considered amines (methylamine, 336 kPa at 298.15 K) (Ge et al., 2011b), it is
- 10 likely that there is also no evaporative loss of amines from solution during the preparation of the aminium sulphate solutions. This result confirms that the estimated concentrations of the $(NH_4)_2SO_4$ solutions prepared from NH_3 and H_2SO_4 are accurate, and that the preparation method is reliable for both $(NH_4)_2SO_4$ and all the six aminium sulphates.

3 Results and Discussion

15 We first report our measurements of the hygroscopic response of the sequence of six aminium salts before comparing our results with previous studies and assessing the accuracy and reproducibility of our data.

3.2 Hygroscopic Properties of Aminium Sulphate Droplets

- The hygroscopic properties of the series of six aminium sulphates were characterised by means of comparative kinetics 20 measurements and using the density and refractive index parameterisations discussed for the CK-EDB data treatment, as described in Sect. 2.1. First, the radial growth curves are shown in Fig. 3a. Considering the compound with the lowest molecular weight first, the hygroscopic behaviour of MMAS is the most similar to ammonium sulphate, in terms of *GFr*. Continuing in the methylaminium sulphates series, a slight decrease in *GF^r* is observed in the high water activity region (*aw*>0.8) for DMAS and TMAS, while at lower water activities, the hygroscopic properties of these compounds converge to 25 (NH4)2SO⁴ within the uncertainties of the measurements as the amount of water in the particles decreases. With respect to the ethylaminium sulphates series, a more evident decreasing trend in the radial growth curve is observed with increasing number of C atoms in the cation (MEAS > DEAS > TEAS), once again especially in the upper part of the curves. If the mono-, diand tri- pairs within the two different series are compared, the methyl compound always presents higher values of radial growth
- 30 presented here and below for TEAS need to be interpreted cautiously because of the uncertainty of the TEA stock solution. However, the trends that have been observed for TEAS when compared to the other five aminium sulphate systems seem to be completely plausible and this may indicate that the assumed initial TEA concentration is reasonable.

factor than its equivalent in the ethylaminium sulphates series. As indicated in Section 2.2, it is worth reiterating that the results

For the calculation of the dry radius reference state in the denominator of *GFr*, the pure melt density is used (Sect. 2.1); if the pure solid density values were known and used, one could expect that the calculated radial growth curve would be slightly higher, a consequence of the solid densities having higher values than the melt density (Clegg et al., 2013). A comparison of

- 5 the pure melt densities (*ρmelt*) from our work and from Clegg et al. (2013) and solid densities (*ρsolid*) estimated by Qiu and Zhang (2012) is provided in Table S4 in the Supplementary Information, with our data previously published in Cai et al. (2016). As an example, if a hypothetical increase of 5% from *ρmelt* to *ρsolid* as a rough estimate is considered for DMAS, the obtained *GF^r* curve would increase by less than 1% (less than 0.01 in *GFr*); this would not affect the trends shown in Fig. 3a, although it would marginally change the relative position of the aminium sulphates curves to that of ammonium sulphate.
- 10

The same datasets shown in Fig. 3a have been converted to *nwater/nsolute* vs. *a^w* and are shown in Fig. 3b. Note that the *y*-axis is displayed on a logarithmic scale for clarity. Different trends can be recognised: the primary factor influencing how many moles of water are absorbed per mole of solute is the number of alkyl groups in the cation and not the length of the carbon chains (i.e. mono- vs. di- vs. tri-methyl or ethyl). In fact, when the *nwater/nsolute* curves (Fig. 3b) are compared with the hygroscopic

- 15 behaviour plotted as *GF^r* (Fig. 3a), more significant differences are apparent in Fig. 3b within each series (e.g. among MMAS, DMAS and TMAS) rather than between analogous compounds belonging to the two series (e.g. between TMAS and TEAS), even though the ethylaminium sulphate in a pair always has slightly higher values of n_{water}/n_{solute} (i.e. MMAS \approx MEAS \lt $DMAS \approx DEAS \leq TMAS \approx TEAS$).
- 20 The seemingly different trends found for the hygroscopic properties of aminium sulphates when reported either as *GF^r* or $n_{\text{water}}/n_{\text{solute}}$ as a function of a_w originate in the different physicochemical properties that each quantity is dependent on. When *GF^r* is calculated, the density and molecular weight of each compound are central to its determination. For example, if particles with a dry radius of 1 µm for each of the six aminium sulphates are considered, the moles of salt in each droplet decreases with increase in molecular weight and decrease in density, as shown in Fig. 4. Therefore, if hygroscopicity is represented by a radial 25 growth factor, *GF_r* is determined not only by the ability of a compound to absorb water at a certain relative humidity, but also
- by the number of solute molecules present in the droplet itself. For this reason, a decreasing hygroscopic growth trend with increasing molecular weight of the aminium sulphates is apparent in Fig. 3a for *GFr*.

Growth factor curves are widely used in the aerosol literature, especially when the optical properties and radiative forcing or 30 the climatic effects of atmospheric aerosols are investigated, because these properties largely depend on particle size. The quantification of the hygroscopic properties of a compound in terms of *GF^r* is therefore useful for such applications and is reported in this work. However, $n_{\text{water}}/n_{\text{solute}}$ vs. a_w curves allow the thermodynamic characterisation of the water uptake of a substance and the decoupling of its hygroscopic properties from its density and molecular weight, with the water uptake 'normalised' to the moles of solute in the particle. This representation – effectively water activity as a function of concentration

– is mostly used when focusing on chemical reactivity or on the thermodynamic properties of compounds in the aerosol phase. The increasing trend apparent in Fig. 3b for *nwater/nsolute* curves with increasing number of carbon atoms in the cation is related to the size of the cation itself: the bigger the cation, the larger the number of water molecules required to solvate the cation at a particular water activity. It is perhaps worth noting here the high level of accuracy in retrieving hygroscopic growth by this 5 method and the opportunity it provides to resolve such fine trends in growth factor.

3.2 Comparisons with Previous Studies

As described in the Introduction, there have been some recent reports of the physicochemical properties of aminium sulphate aerosols and aqueous solutions, motivated by the increasing understanding of their role in atmospheric processes (Bzdek et al.,

- 10 2010; Ge et al., 2011a; Lavi et al., 2015; Liu et al., 2012; Qiu and Zhang, 2013). With respect to the determination of their hygroscopic properties, the approach presented here differs from the earlier reports in either the preparation method of the solutions or in the experimental measurement technique. We now compare our measurements using a CK-EDB with the results of these previous studies.
- 15 Qiu and Zhang (2012) were the first to measure diameter growth curves for these compounds (excluding MEAS) by means of an HTDMA (Hygroscopicity Tandem Differential Mobility Analyser). They inferred hygroscopic growth factors from the ratio of the mobility diameters measured at a variable RH (up to 90%) and at an RH of ∼12%. A monotonic increase in the particle size with RH for each compound was observed and no deliquescence/efflorescence behaviour was observed. If their growth curves are compared with the *GF^r* shown in Fig. 3a (Fig. S3 in Supplementary Information), a systematically smaller growth 20 factor is reported from the HTDMA measurements. The largest deviations are for TEAS (-16% on average in the *a^w* range where the two datasets overlap, from ∼ 0.5 to 0.9) and for TMAS (-13%), a deviation of -10% is observed for DEAS and DMAS and of -7% for MMAS. These discrepancies are most likely due to the presence of some residual water at the conditions at which the reference diameter was measured in the HTDMA experiments (RH ∼12%), which would result in an overestimated reference 'dry' size and in underestimated growth factor values. This explanation is supported by the studies of 25 Chan and Chan (2012) who reported the presence of water for some aminium sulphates even at an RH of ∼3%. In addition, the possible volatilisation of the amine during the drying step in the HTDMA would alter the chemical composition of the particles (Chan and Chan, 2012, 2013), not only artificially reducing the apparent dry size recorded but shifting the dry particle
	- towards an aminium bisulphate composition (1:1 molar ratio of sulphuric acid to amine), which has been shown to be less hygroscopic than its sulphate counterpart (Sauerwein et al., 2015).
- 30

Consistent with the above discussion, the two compounds with the biggest deviation between the *GF^r* values reported here and by Qiu and Zhang (TMAS and TEAS) are those that were found to be affected by the largest evaporative losses of amine by Chan and Chan (2012). They reported studies of ammonium displacement reactions by alkylamines by levitating single droplets in an EDB and evaluating the changes in the recorded Raman spectra during an experiment. After TEAS solution droplets were levitated at RH<3% for more than 5 h, TEAS was found to have converted to TEA bisulphate almost completely, indicating that half of the amine in the initial solution droplet had evaporated. Similar behaviour was observed for TMAS, while DMAS and DEAS showed a smaller degree of evaporation of the amine (∼25% and ∼5%, respectively); MMAS and

- 5 MEAS did not show any relevant evaporation over the experimental timescales. These results support the hypothesis of possible evaporation of the amine from the drying solution droplets during HTDMA experiments, and at the same time represent an interesting comparison to evaluate the timescales over which the evaporative loss of amine is significant. The hygroscopic growth measurement of each droplet of aminium sulphate solution in CK-EDB studies extends over 30 s at the longest: this allows the effective decoupling of the fast water evaporation and the slow amine evaporation, since these two
- 10 processes occur over different timescales. In addition, the hygroscopicity measurements in the present work were carried out at RHs in excess of 50% at all times; for this reason, the evaporation of the amines are necessarily smaller than those measured by Chan and Chan (2012).

Clegg et al. (2013) have converted the size growth curves measured by Qiu and Zhang to $n_{\text{water}}/n_{\text{solute}}$ vs. a_w . The growth curves 15 that resulted did not show any discernible trends in the hygroscopicity of aminium sulphates. Indeed, the results were essentially similar to the water uptake of ammonium sulphate in the lower a_w range. At higher a_w above 0.7, the scatter in the data limited the identification of any apparent trend apart from the suggestion that the methylaminium sulphates absorb fewer moles of water per moles of solute than the compounds in the ethyl series. Fig. 5 shows the large differences between the water uptake determined from the results of Qiu and Zhang, and that from the work in this study. Because of the uncertainties 20 associated with the HTDMA *GF^r* values, it is hard to draw conclusions the comparison with our results.

Sauerwein et al. (2015) have recently reported bulk water activity measurements with an activity meter (Aqua lab Series 3TE) for various amine-to-sulphate ratios and over a concentration range of dissolved electrolyte up to 9 mol $kg⁻¹$ at 25 $^{\circ}$ C. (Note that TEAS was not considered in their study.) Using a bulk measurement technique for the determination of hygroscopic

- 25 properties of aminium sulphates has the advantage of limiting possible evaporative losses of the amine during the experiments, especially compared to accumulation-mode aerosol measurements with an HTDMA for which timescales of evaporative loss/equilibration are assumed very short. As a first comparison, the hygroscopic growth curves previously shown in Fig. 3b in terms of *nwater/nsolute* vs. *a^w* are compared with results from Sauerwein et al. (2015) (Fig. 6a in their paper) in Fig. 5. The hygroscopic growth curves determined from the two different experimental methods differ marginally, but these discrepancies
- 30 are considerably smaller than the differences between the results of both studies CK-EDB and bulk water activity measurements – and HTDMA data from Clegg et al. (2013).

A comparison of the two datasets represented in terms of osmotic coefficients (Eq. (2)) can provide further insight, see Fig. 6. Notably, the measurements reported here extend to both more dilute and concentrated solutions and include measurements for

TEAS. If the errors associated with each dataset are considered, the two sets of measurements overlap over some of the range in *m*(SO₄^{2-)0.5}, especially in the low molality limit (high water activity region) for MMAS, TMAS, MEAS and DEAS (*m*(SO₄²⁻ $0.5 < 1$, approximately), but also in the region $m(SO_4^2)^{0.5} > 2$ for DMAS and TMAS. A comparison with the uncertainty in the osmotic coefficients that arises from an uncertainty of ± 0.002 in water activity (same as in Fig. 2b) indicates that the observed

- 5 discrepancies are larger in magnitude than this typical experimental uncertainty of our technique, except for very low sulphate molality values (below ~ 0.8). However, despite the discrepancies that do exist, results from the two different experimental methods show the same qualitative hygroscopicity trends over the methyl and ethylaminium sulphates series. For further comparison, the results of Clegg et al. (2013) (from the HTDMA data) were converted to stoichiometric coefficients and are plotted in Fig. 6. These values are broadly comparable for all of the compounds, and close to those for ammonium sulphate as
- 10 noted earlier. They do not agree with either the results of the experiments reported here, or the bulk measurements of Sauerwein et al. (2015).

We now consider the differences that exist between the data of Sauerwein et al. (2015) and our measurements. It is worth stating here that the evaporation kinetics measurements at the core of the CK-EDB approach have been validated in a previous 15 publication (Rovelli et al., 2016) by means of the determination of the hygroscopic properties of well-characterised inorganic compounds and their mixtures. In addition, we have performed sensitivity tests to evaluate possible effects of random experimental errors associated with the proposed experimental method. The reproducibility of our measurements and the uncertainties associated with the treatment of densities are discussed below in Sect. 3.3 and 3.4, where they are shown not to

affect significantly the determined hygroscopic properties of a compound.

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If the differences between our measurements and the results of Sauerwein et al. (2015) were caused by partial volatilisation of the amine from the droplets evaporating in the CK-EDB, the observed bias between the two datasets would be reversed, i.e. an underestimation of *nwater/nsolute* and of osmotic coefficients would be expected, similar to what has been discussed in the case of the HTDMA size-based measurements of Qiu and Zhang (2012) when converted to a molar basis by Clegg et al. (2013).

- 25 Another potential significant source of error could be inaccuracies in the starting concentrations of the aminium sulphate solutions. In this respect, the solution preparation and measurement method described in Sect. 2.2 has been validated through the reaction of ammonia and sulphuric acid and the measurements of the hygroscopic properties of the obtained $(NH_4)_{2}SO_4$. If the volatilisation of NH₃ during the neutralisation reaction was not an issue in preparing the solution of $(NH₄)₂SO₄$ it is unlikely to present a problem for the alkylamine solutions, with all alkylamines characterised by lower vapour pressures than ammonia
- 30 (Ge et al., 2011b). Furthermore, in order to achieve complete agreement with the data from Sauerwein et al. (2015) the aminium sulphates solutions should be 5-10% more concentrated than calculated from the titration of the commercial stock solutions and from the amount of the reagents. Not only does this percentage seem unrealistically high, but the solute concentrations would be overestimated and not underestimated if any volatilisation of the amine occurred during the preparation of the solutions.

Sauerwein et al. (2015) provide a framework for the estimation of the water content in mixtures (in this case ternary mixtures of H2O-H2SO4-amine) at any molar ratio of solutes. This allows a calculation of the water content (in terms of kg of water per mole of solute, or molality) and the stoichiometric osmotic coefficients (Eq. (2)) for the relevant aminium:sulphate ratio of 1:1

- 5 (bisulphate) and 2:1 (sulphate). However, it should be noted that unlike the study presented here, measurements were not explicitly made for the 2:1 molar ratio aqueous aminium sulphate solution. Instead, Sauerwein et al. used a modified Zdanovskii-Stokes-Robinson expression (ZSR, Eq. (3) of Sauerwein et al.) to represent the water uptake of the different mixtures of aqueous aminium sulphates and H2SO4; these fits are shown in Fig. 7 along with the actual measured points. *xsalt* can be considered as the degree of neutralisation of sulphuric acid. *xsalt*=0 corresponds to pure H2SO⁴ and the water content in
- 10 this case is well-known (Clegg and Brimblecombe, 1995); the ZSR fit is therefore constrained to this value. A value at *xsalt*=1 corresponds to the completely neutralised aminium sulphate. The results of the fit were used to estimate the water uptake of each solution of aqueous aminium sulphate (i.e. the 2:1 molar ratio at $x_{\text{val}}=1$) and it is these values that are shown in Fig. 6. In order to compare our data with that of Sauerwein et al., values at the same water activities must be compared (0.8 and 0.925 are chosen here) and a linear interpolation between actual measurement points is required to achieve this. Normally, the data
- 15 from CK-EDB measurement are calculated with a_w steps of 0.01 for $a_w > 0.8$ where hygroscopic growth curves are the steepest, with a spacing of 0.02 for the water activity range ~ 0.65 - 0.8 and of 0.03 for $a_w < 0.65$. As in Sauerwein et al. (2015), the uncertainty of the interpolated points is set as the largest uncertainty among the experimental points used for the interpolation. The discrepancies between the estimations from bulk and CK-EDB measurements shown in Fig. 5 are apparent in Fig. 7 too at $x_{\text{sal}}=1$: the amount of water associated with each mole of aminium sulphate is systematically higher from the CK-EDB
- 20 measurements when compared with the bulk measurements at both *a^w* with the discrepancy higher at the higher water activity.

In view of the differences between our measurements at $x_{\text{salt}} = 1.0$ and the extrapolations from the fits of Sauerwein et al., we include a second fit to the ZSR model used previously but including our own measurements of the water uptake. The fit equations and parameters are provided in Supplementary Information (Table S6). As well as the fitted values more closely 25 reflecting our measured values for the pure aminium sulphates at *x*salt equal to unity, the revised fit does not significantly worsen the fit to the measurements of Sauerwein et al. for the acidified mixtures when *x*salt < 0.6). We therefore suggest that the discrepancies seen in terms of $n_{\text{water}}/n_{\text{solute}}$ (Fig. 5), osmotic coefficients (Fig. 6) and kg of water per mole of solute in Fig. 7 can be attributed, in part, to the uncertainties associated with the ZSR fits and interpolation necessary to estimate the water

30 why a simple relation such as the modified ZSR equation (Eq. (3) of Sauerwein et al.) should exactly represent the relationship between water activity and chemical composition in such systems, especially considering the large and unquantified influence of the sulphate-bisulphate equilibrium on the thermodynamic properties of the solutions.

uptake of the pure aqueous aminium sulphates from the results of Sauerwein et al. (2015). Certainly there is no physical reason

A possible constant error in measured water activity could be the origin of the differences, but this seems unlikely. Such an offset in water activity results in a large change in the osmotic coefficient at the lowest concentrations for which measurements were made, but a much smaller change at the highest concentrations. This is not consistent with the differences in ϕ_y shown in Fig. 6. As a final consideration, we now explore the reproducibility of the CK-EDB measurements and the sensitivity of the 5 retrieved hygroscopicity to the chosen treatment for solution density.

3.3 Reproducibility of Measurements

A full validation of the retrieval of the hygroscopic properties of single trapped solution droplets from CK-EDB experiments has already been presented in a previous publication (Rovelli et al. 2016), where we demonstrated the accuracy of the approach 10 by reporting hygroscopicity measurements for well-characterised inorganic components. In this Section, we evaluate the reproducibility of the solution preparation method (Sect. 2.2) with the aim of demonstrating that the results presented in the previous Sections are not affected by any random error associated with our approachfor making the sample solutions. Data obtained from three different datasets of evaporating DMAS solution droplets are compared in Fig. 8. The variation in $n_{\text{water}}/n_{\text{solute}}$ with a_w (Panel (a)) and the change in osmotic coefficient with square root of sulphate mass fraction (Panel (b)) are

- 15 calculated from three different data sets, each arising from measurements with 10 droplets and prepared from different starting stock solution of diethylaminium sulphate, obtained separately one from the other by mixing DEA and H_2SO_4 in different days. Data from Sauerwein et al. (2015) (open circles) and calculations for ammonium sulphate (E-AIM model, line) are shown for comparison. Measurements were made with droplet evaporation into a gas phase RH of ~80%. The reproducibility of the data is very satisfactory, giving further evidence that the applied solution preparation procedure coupled to the retrieval of the
- 20 hygroscopic properties with CK-EDB experiments is reliable. In addition, this consistency in experimental reproducibility strongly suggests that the discrepancies with the data of Sauerwein et al. (2015) do not originate from random errors associated with the CK-EDB experiments.

3.4 Sensitivity to Parameterization of Solution Density

- 25 A knowledge of the solution density as a function of solute concentration is needed to the process CK-EDB evaporation radius profiles (Sect. 2.1). The effect of uncertainties in the density parameterization used is here evaluated for DMAS, in order to estimate how potential errors in the density measurements and uncertainties in the extrapolation of ρ_{melt} at *mfs* equal to unity could affect the hygroscopicity data retrieved from comparative kinetics measurements. In Fig. 9, the densities of bulk DMAS solutions measured in this work are shown together with their 3rd order polynomial fit, which is the parameterization used in
- 30 all the calculations for this compound in this work (Table 1). As shown in a previous work (Cai et al., 2016), a $\pm 2\%$ uncertainty is typical for the estimated melt density ($mfs = 1.0$) predicted from with a 3rd order polynomial fit of measured solution densities (Sect. 2.1), while the error is decreases with $mfs^{0.5}$ for the rest of the curve as shown in Fig. 9. More generally, the aminium

sulphates are very soluble compounds and it was possible to directly measure the density of their bulk solutions over a wide range of solute mass fractions (up to *mfs0.5* values of 0.67-0.85, depending on the compound, see Table S2). Consequently, the *mfs*^{0.5} range over which the extrapolation for the calculation of ρ_{melt} is needed is small and an uncertainty of $\pm 2\%$ (Fig. 9) is large for the evaluation of its effects. The two 3rd order polynomials (called 'Error+' and 'Error-' in Table 1, respectively) were

- 5 applied together with the molar refraction mixing rule to explore the sensitivity of the data analysis of the original DMAS evaporation kinetics datasets to the representation of the solution density. Density data from Clegg et al. (2013) are also shown in Fig. 9 for comparison. A small discrepancy between their measurements and the data presented in this work is observed, not only for DMAS but also for the other 4 aminium sulphates they considered; but no clear pattern to these differences. Figure S4 shows a further comparison of density data from this work and from Clegg et al. (2013) converted in terms of apparent
- 10 molar volumes; both measured and fitted apparent molar values are provided and the coefficient of the fitted equations can be found in Table S5. They may be attributable to the different preparation procedures of the aminium sulphates solutions and to the different experimental techniques for the measurement of densities.
- The hygroscopic properties of DMAS obtained from the treatment of the evaporation kinetics data with the 'original' density 15 parameterization and with the upper and lower bounds on the density treatment are shown in Fig. 10. Data from Sauerwein et al. (2015) and calculations for ammonium sulphate (E-AIM model) are plotted for comparison. When the hygroscopic properties of DMAS are represented either in terms of the dependence of either *nwater/nsolute* or *GF^r* on water activity (Panels (a) and (b)), the three curves deriving from the three different density treatments are virtually undistinguishable; thus the uncertainty in the applied density parameterisation does not significantly alter the analysis. In the case of the osmotic 20 coefficients plot (Panel (c)) some very slight deviations between the three treatments can be distinguished at the two extremes of the plot. With respect to the low sulphate molality region, these small differences are due to the fact that a small variation in the sulphate molality results in more significant variations in the osmotic coefficients, because *mst* appears in the denominator in the osmotic coefficient expression (Eq. (2)). If the high sulphate molality region is considered, the variations among the three curves are more significant because the simulated error on the density parameterisation is larger for more concentrated
- 25 solutions (i.e. the size of the grey envelope in Fig. 9).

It is clear that the hygroscopic properties of aminium sulphates determined from CK-EDB measurements are relatively insensitive to reasonable variations in the extrapolated ρ_{melt} value ($\pm 2\%$) and to the applied density parameterisation. In addition, the variations introduced by different density parameterisations are very small compared to the differences from the

30 results of Sauerwein et al., who obtain lower values of the hygroscopicity and osmotic coefficient. We conclude that these differences cannot be caused by inaccuracies in the approaches for treating variations of the density of evaporating droplets.

4 Atmospheric Importance and Conclusions

Quantifying the hygroscopic properties of aminium sulphates is important for understanding and modelling of the atmospheric processes in which they are involved. In particular, the role of short-chained alkylamines in the formation of new particles has been investigated in recent literature studies and found to be significant (Section 1). Aminium sulphate-rich nanoparticles that

- 5 derive from new particles formation events can potentially act as CCN, and their hygroscopic properties must be wellcharacterised with the aim of reducing the overall uncertainties that currently affect our understanding of the indirect effects of atmospheric aerosols on climate. Robust and accurate data are essential for improving microphysical models of aerosol hygroscopicity; this study presents an extensive data set for an homologous series of six compounds, compared to ammonium sulphate, extending over a wide range in RH. In addition, it represents the most comprehensive characterisation of the
- 10 hygroscopic response of aminium sulphate aerosol so far, complementing previous bulk phase measurements (comparable in accuracy but limited to higher water activity) and aerosol measurements at lower RH (with lower accuracy than achieved here). Previously, the bulk and aerosol measurements reported in the literature were in disagreement. Here, we report aerosol measurements that are in good agreement with the previously most accurate bulk phase data, resolving this discrepancy.
- 15 In order to measure their hygroscopic properties with a CK-EDB, the dependence of aminium sulphates solution densities and refractive indices on mass fraction of solute are reported. The experimental technique together with the application of the molar refractive mixing rule and a 3rd order polynomial parameterisation of density for the representation of the refractive indices and densities of solutions with variable solute mass fractions were presented and validated in previous works (Cai et al., 2016; Davies et al., 2013; Rovelli et al., 2016).The procedure for preparation of the aminium sulphates stock solutions was validated 20 by using it for the preparation of ammonium sulphate solutions from the direct reaction of ammonia and sulphuric acid. The hygroscopic properties of the obtained (NH4)2SO⁴ solutions retrieved from CK-EDB measurements were in very good agreement with calculations from E-AIM model and we demonstrated that the uncertainty of such measured values was comparable with a typical uncertainty of ± 0.002 on a_w , which was previously shown to be typical for CK-EDB measurements (Rovelli et al., 2016). This result demonstrated that the used preparation procedure is robust and reliable.

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The experimental results for the aminium sulphates were compared with the few studies available in the literature and the observed discrepancies were discussed in the light of the different experimental approaches. Measurements from our new approach provide a level of accuracy that reveals clearly the fine variations in hygroscopic growth (down to discernible difference is *GF_r* of order 0.01-0.02) that occur with molecular structure and substitution, and avoids the additional complexity

30 of volatilisation of semi-volatile components during hygroscopic growth with measurements complete in a matter of a few seconds. The largest discrepancies with previous data were found when comparing the CK-EDB results with H-TDMA measurements (Clegg et al., 2013; Qiu and Zhang, 2012). These differences can be attributed to a possible overestimation of the dry size of particles due to residual water in the reference dry state in HTDMA measurements and/or to a shift of the chemical composition of particles towards the bisulphate composition because of partial volatilization of the amines from solution. Closer comparability was found with the bulk water activity measurements by Sauerwein et al. (2015). The main differences in approaches are that: we perform aerosol measurements that cover a wider range in water activity as compared with the bulk measurements of Sauerwein et al. (2015); and we provide direct measurement at amine-to-sulphate ratios of

- 5 exactly 2:1, whereas Sauerwein et al. (2015) performed a ZSR fitting on data from solutions with variable amine-to-sulphates ratios and extrapolated water content for the exact 2:1 ratio. These new CK-EDB measurements suggest a higher level of hygroscopic growth for the aminium sulphates than previously reported by Sauerwein and co-workers when inferred from measurements over a range of amine-to-sulphates ratios; we have provided a refined parameterisation for all compositions.
- 10 It is worth noting that the characterization of the hygroscopic properties of aminium sulphates up to *a^w* of 0.99 was possible with the CK-EDB technique. The other literature approaches that were discussed in this Section were able to cover larger *a^w* ranges (down to 0.1) but none of them could be applied to obtain any data for $a_w > 0.9$. Thus, the comparative kinetics measurements in a CK-EDB provide a powerful tool for investigating a water activity region that is otherwise hard to characterise with such accuracy, but which is of great importance for the understanding of the activity of aerosol particles as
- 15 cloud condensation nuclei (Wex et al., 2009).

As a final remark, aminium sulphates are the first class of mixed inorganic-organic aerosol systems to be investigated by means of the CK-EDB comparative kinetics technique, described in Rovelli et al. (2016), over a wide range of water activities. Therefore, besides the atmospheric relevance of these compounds, this study also provides a deeper understanding of the

20 possible effects caused by random errors in the experimental procedure, and by uncertainties on the representation of the density of a compound. We demonstrated that the CK-EDB measurements are characterised by a very good level of reproducibility and that a typical ±2% uncertainty on the value of the extrapolated melt density only marginally affect the measured hygroscopic properties.

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References

Akyüz, M.: Simultaneous determination of aliphatic and aromatic amines in indoor and outdoor air samples by gas chromatography-mass spectrometry, Atmos. Environ., 42, 3809–3819, doi:10.1016/j.talanta.2006.10.028, 2008.

Albrecht, B. A.: Aerosols, Cloud microphysics, and Fractional Cloudiness, Science (80-.)., 245, 1227–1230, 1989.

- 5 Almeida, J., Schobesberger, S., Kürten, A., Ortega, I. K., Kupiainen-Määttä, O., Praplan, A. P., Adamov, A., Amorim, A., Bianchi, F., Breitenlechner, M., David, A., Dommen, J., Donahue, N. M., Downard, A., Dunne, E., Duplissy, J., Ehrhart, S., Flagan, R. C., Franchin, A., Guida, R., Hakala, J., Hansel, A., Heinritzi, M., Henschel, H., Jokinen, T., Junninen, H., Kajos, M., Kangasluoma, J., Keskinen, H., Kupc, A., Kurtén, T., Kvashin, A. N., Laaksonen, A., Lehtipalo, K., Leiminger, M., Leppä, J., Loukonen, V., Makhmutov, V., Mathot, S., McGrath, M. J., Nieminen, T., Olenius, T., Onnela, A., Petäjä, T., Riccobono,
- 10 F., Riipinen, I., Rissanen, M., Rondo, L., Ruuskanen, T., Santos, F. D., Sarnela, N., Schallhart, S., Schnitzhofer, R., Seinfeld, J. H., Simon, M., Sipilä, M., Stozhkov, Y., Stratmann, F., Tomé, A., Tröstl, J., Tsagkogeorgas, G., Vaattovaara, P., Viisanen, Y., Virtanen, A., Vrtala, A., Wagner, P. E., Weingartner, E., Wex, H., Williamson, C., Wimmer, D., Ye, P., Yli-Juuti, T., Carslaw, K. S., Kulmala, M., Curtius, J., Baltensperger, U., Worsnop, D. R., Vehkamäki, H. and Kirkby, J.: Molecular understanding of sulphuric acid-amine particle nucleation in the atmosphere., Nature, 502, 359–363, doi:10.1038/nature12663,
- 15 2013.

Barnes, M. D., Lermer, N., Whitten, W. B. and Ramsey, J. M.: A CCD based approach to high-precision size and refractive index determination of levitated microdroplets using Fraunhofer diffraction, Rev. Sci. Instrum., 68(6), 2287, doi:10.1063/1.1148134, 1997.

Barsanti, K. C., McMurry, P. H. and Smith, J. N.: The potential contribution of organic salts to new particle growth, Atmos. 20 Chem. Phys., 9, 2949–2957, doi:10.5194/acp-9-2949-2009, 2009.

Broday, D. M. and Georgopoulos, P. G.: Growth and Deposition of Hygroscopic Particulate Matter in the Human Lungs, Aerosol Sci. Technol., 34(34), 144–159, doi:10.1080/027868201300082184, 2001.

Bzdek, B. R., Ridge, D. P. and Johnston, M. V.: Amine exchange into ammonium bisulfate and ammonium nitrate nuclei, Atmos. Chem. Phys. Discuss., 10(1), 45–68, doi:10.5194/acpd-10-45-2010, 2010.

25 Cai, C., Miles, R. E. H., Cotterell, M. I., Marsh, A., Rovelli, G., Rickards, A. M. J., Zhang, Y.-H. and Reid, J. P.: Comparison of methods for predicting the composition dependence of the density and refractive index of organic-aqueous aerosols, J. Phys. Chem. A, 120, 6604–6617, doi:10.1021/acs.jpca.6b05986, 2016.

Chan, L. P. and Chan, C. K.: Displacement of Ammonium from Aerosol Particles by Uptake of Triethylamine, Aerosol Sci. Technol., 46(2), 236–247, doi:10.1080/02786826.2011.618815, 2012.

30 Chan, L. P. and Chan, C. K.: Role of the aerosol phase state in ammonia/amines exchange reactions, Environ. Sci. Technol., 47(11), 5755–5762, doi:10.1021/es4004685, 2013.

Choi, M. Y. and Chan, C. K.: The effects of organic species on the hygroscopic behaviors of inorganic aerosols, Environ. Sci. Technol., 36(11), 2422–2428, doi:10.1021/es0113293, 2002.

Chu, Y., Sauerwein, M. and Chan, C. K.: Hygroscopic and phase transition properties of alkyl aminium sulfates at low relative humidities, Phys. Chem. Chem. Phys., 17(30), 19789–19796, doi:10.1039/C5CP02404H, 2015.

Clegg, S. L. and Brimblecombe, P.: Application of a Multicomponent Thermodynamic Model to Activities and Thermal Properties of 0-40 mol kg^{\wedge -1 Aqueous Sulfuric Acid from <200 to 328 K, J. Chem. Eng. Data, 40(1), 43–64,} 5 doi:10.1021/je00017a012, 1995.

- Clegg, S. L., Qiu, C. and Zhang, R.: The deliquescence behaviour, solubilities, and densities of aqueous solutions of five methyl- and ethyl-aminium sulphate salts, Atmos. Environ., 73, 145–158, doi:10.1016/j.atmosenv.2013.02.036, 2013. Davies, J. F.: Characterising The Impact Of Bulk, Surface And Gas-Phase Limitations On Mass Transport In Aerosol., 2014. Davies, J. F., Haddrell, A. E., Miles, R. E. H., Bull, C. R. and Reid, J. P.: Bulk, surface, and gas-phase limited water transport
- 10 in aerosol, J. Phys. Chem. A, 116(45), 10987–10998, doi:10.1021/jp3086667, 2012a. Davies, J. F., Haddrell, A. E. and Reid, J. P.: Time-Resolved Measurements of the Evaporation of Volatile Components from Single Aerosol Droplets, Aerosol Sci. Technol., 46(6), 666–677, doi:10.1080/02786826.2011.652750, 2012b. Davies, J. F., Haddrell, A. E., Rickards, A. M. J. and Reid, J. P.: Simultaneous analysis of the equilibrium hygroscopicity and water transport kinetics of liquid aerosol, Anal. Chem., 85(12), 5819–5826, doi:10.1021/ac4005502, 2013.
- 15 Dennis-Smither, B. J., Marshall, F. H., Miles, R. E. H., Preston, T. C. and Reid, J. P.: Volatility and oxidative aging of aqueous maleic acid aerosol droplets and the dependence on relative humidity, J. Phys. Chem. A, 118(30), 5680–5691, doi:10.1021/jp504823j, 2014.

DePalma, J. W., Bzdek, B. R., Doren, D. J. and Johnston, M. V: Structure and Energetics of Nanometer Size Clusters of Sulfuric Acid with Ammonia and Dimethylamine, J. Phys. Chem. A, 116(3), 1030–1040, doi:10.1021/jp210127w, 2012.

20 Dusek, U., Frank, G. P., Hildebrandt, L., Curtius, J., Schneider, J., Walter, S., Chand, D., Drewnick, F., Hings, S., Jung, D., Borrmann, S. and Andreae, M. O.: Size matters more than chemistry for cloud-nucleating ability of aerosol particles., Science, 312(5778), 1375–8, doi:10.1126/science.1125261, 2006.

Dutcher, C. S., Ge, X., Wexler, A. S. and Clegg, S. L.: An Isotherm-Based Thermodynamic Model of Multicomponent Aqueous Solutions, Applicable Over the Entire Concentration Range, J. Phys. Chem. A, 117, 3198−3213, 2013.

25 Ge, X., Wexler, A. S. and Clegg, S. L.: Atmospheric amines - Part I. A review, Atmos. Environ., 45(3), 524–546, doi:10.1016/j.atmosenv.2010.10.012, 2011a. Ge, X., Wexler, A. S. and Clegg, S. L.: Atmospheric amines - Part II. Thermodynamic properties and gas/particle partitioning, Atmos. Environ., 45(3), 561–577, doi:10.1016/j.atmosenv.2010.10.013, 2011b.

Glantschnig, W. J. and Chen, S. H.: Light scattering from water droplets in the geometrical optics approximation., Appl. Opt., 30 20(14), 2499–2509, doi:10.1364/AO.20.002499, 1981.

Haddrell, A. E., Davies, J. F., Yabushita, A. and Reid, J. P.: Accounting for changes in particle charge, dry mass and composition occurring during studies of single levitated particles, J. Phys. Chem. A, 116(40), 9941–9953, doi:10.1021/jp304920x, 2012.

Haddrell, A. E., Hargreaves, G., Davies, J. F. and Reid, J. P.: Control over hygroscopic growth of saline aqueous aerosol using

Pluronic polymer additives, Int. J. Pharm., 443(1–2), 183–192, doi:10.1016/j.ijpharm.2012.12.039, 2013.

Haddrell, A. E., Davies, J. F. and Reid, J. P.: Dynamics of Particle Size on Inhalation of Environmental Aerosol and Impact on Deposition Fraction, Environ. Sci. Technol., 49(24), 14512–14521, doi:10.1021/acs.est.5b01930, 2015.

Heinisch, C., Wills, J. B., Reid, J. P., Tschudi, T. and Tropea, C.: Temperature measurement of single evaporating water 5 droplets in a nitrogen flow using spontaneous Raman scattering., Phys. Chem. Chem. Phys., 11(42), 9720–9728, doi:10.1039/b908555f, 2009.

Kallinger, G. and Niessner, R.: Laboratory investigation of annular denuders as sampling system for the determination of aliphatic primary and secondary amines in stack gas, Mikrochim. Acta, 130(4), 309–316, 1999.

Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C., Van Dingenen, R., Ervens, B., Nenes,

T., Hanson, D. R., Kulmala, M., McFiggans, G., Pöschl, U., Riipinen, I., Rossi, M. J., Rudich, Y., Wagner, P. E., Winkler, P.

- 10 A., Nielsen, C. J., Swietlicki, E., Putaud, J. P., Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L., Tsigaridis, K., Vignati, E., Stephanou, E. G. and Wilson, J.: Organic aerosol and global climate modelling: a review, Atmos. Chem. Phys., 5, 1053–1123, doi:10.5194/acp-5-1053-2005, 2005. Kolb, C. E., Cox, R. A., Abbatt, J. P. D., Ammann, M., Davis, E. J., Donaldson, D. J., Garrett, B. C., George, C., Griffiths, P.
- 15 M., Worsnop, D. R. and O' Dowd, C. D.: An overview of current issues in the uptake of atmospheric trace gases by aerosols and clouds, Atmos. Chem. Phys., 10(21), 10561–10605, doi:10.5194/acp-10-10561-2010, 2010. Kulmala, M., Vesala, T. and Wagner, P. E.: An Analytical Expression For the Rate of Binary Condensational Particle Growth, Proc. R. Soc. A Math. Phys. Eng. Sci., 441(1913), 589–605, doi:10.1098/rspa.1993.0081, 1993. Kulmala, M., Kontkanen, J., Junninen, H., Lehtipalo, K., Manninen, H. E., Nieminen, T., Petäjä, T., Sipilä, M., Schobesberger,
- 20 S., Rantala, P., Franchin, A., Jokinen, T., Järvinen, E., Äijälä, M., Kangasluoma, J., Hakala, J., Aalto, P. P., Paasonen, P., Mikkilä, J., Vanhanen, J., Aalto, J., Hakola, H., Makkonen, U., Ruuskanen, T., Mauldin, R. L., Duplissy, J., Vehkamäki, H., Bäck, J., Kortelainen, A., Riipinen, I., Kurtén, T., Johnston, M. V, Smith, J. N., Ehn, M., Mentel, T. F., Lehtinen, K. E. J., Laaksonen, A., Kerminen, V.-M. and Worsnop, D. R.: Direct observations of atmospheric aerosol nucleation., Science, 339(6122), 943–6, doi:10.1126/science.1227385, 2013.
- 25 Lavi, A., Bluvshtein, N., Segre, E., Segev, L., Flores, M. and Rudich, Y.: Thermochemical, cloud condensation nucleation ability, and optical properties of alkyl aminium sulfate aerosols, J. Phys. Chem. C, 117(43), 22412–22421, doi:10.1021/jp403180s, 2013.

Lavi, A., Segre, E., Gomez-Hernandez, M., Zhang, R. and Rudich, Y.: Volatility of Atmospherically Relevant Alkylaminium Carboxylate Salts, J. Phys. Chem. A, 150126161336007, doi:10.1021/jp507320v, 2015.

30 Lee, J., Carrascón, V., Gallimore, P., Fuller, S., Björkegren, A., Spring, D., Pope, F. and Kalberer, M.: The effect of humidity on the ozonolysis of unsaturated compounds in aerosol particles., Pccp, 14(22), 8023–31, doi:10.1039/c2cp24094g, 2012. Lide, D. R.: CRC Handbook of Chemistry and Physics, 85th ed., CRC Press., 2005. Liu, Y., Ma, Q. and He, H.: Heterogeneous uptake of amines by citric Acid and humic Acid., Environ. Sci. Technol., 46(20), 11112–8, doi:10.1021/es302414v, 2012.

Loukonen, V., Kurtén, T., Ortega, I. K., Vehkamäki, H., Pádua, a. a. H., Sellegri, K. and Kulmala, M.: Enhancing effect of dimethylamine in sulfuric acid nucleation in the presence of water – a computational study, Atmos. Chem. Phys. Discuss., 10(2), 2321–2356, doi:10.5194/acpd-10-2321-2010, 2010.

Mäkelä, J., Yli‐Koivisto, S. and Hiltunen, V.: Chemical composition of aerosol during particle formation events in boreal 5 forest, Tellus B, 380–393, doi:10.1034/j.1600-0889.2001.d01-27.x, 2001.

- McFiggans, G., Artaxo, P., Baltensperger, U., Coe, H., Facchini, M. C., Feingold, G., Fuzzi, S., Gysel, M., Laaksonen, A., Lohmann, U., Mentel, T. F., Murphy, D. M., O'Dowd, C. D., Snider, J. R. and Weingartner, E.: The effect of physical and chemical aerosol properties on warm cloud droplet activation, Atmos. Chem. Phys. Discuss., 5(5), 8507–8646, doi:10.5194/acpd-5-8507-2005, 2005.
- 10 Merikanto, J., Spracklen, D. V., Mann, G. W., Pickering, S. J. and Carslaw, K. S.: Impact of nucleation on global CCN, Atmos. Chem. Phys. Discuss., 9, 8601–8616, doi:10.5194/acpd-9-12999-2009, 2009. Miles, R. E. H., Reid, J. P. and Riipinen, I.: Comparison of approaches for measuring the mass accommodation coefficient for the condensation of water and sensitivities to uncertainties in thermophysical properties, J. Phys. Chem. A, 116(44), 10810– 10825, doi:10.1021/jp3083858, 2012.
- 15 Namieśnik, J., Jastrzębska, A. and Zygmunt, B.: Determination of volatile aliphatic amines in air by solid-phase microextraction coupled with gas chromatography with flame ionization detection, J. Chromatogr. A, 1016(1), 1–9, doi:10.1016/S0021-9673(03)01296-2, 2003.

Ortega, I. K., Kupiainen, O., Kurtén, T., Olenius, T., Wilkman, O., McGrath, M. J., Loukonen, V. and Vehkamäki, H.: From quantum chemical formation free energies to evaporation rates, Atmos. Chem. Phys., 12(1), 225–235, doi:10.5194/acp-12- 20 225-2012, 2012.

Peng, C. G., Chow, A. H. L. and Chan, C. K.: Study of the hygroscopic properties of selected pharmaceutical aerosols using single particle levitation, Pharm. Res., 17(9), 1104–1109, doi:10.1023/A:1026409813779, 2000. Peng, C. G., Chow, A. H. L. and Chan, C. K.: Hygroscopic study of glucose, citric acid, and sorbitol using an electrodynamic balance: Comparison with UNIFAC predictions, Aerosol Sci. Tech., 35(3), 753–758, doi:10.1080/02786820152546798, 2001.

- 25 Petters, M. D. and Kreidenweis, S. M.: A single parameter representation of hygroscopic growth and cloud condensation nucleus activity, Atmos. Chem. Phys., 7, 1961–1971, doi:10.5194/acp-7-1961-2007, 2007. Pope, F. D., Dennis-Smither, B. J., Griffiths, P. T., Clegg, S. L. and Cox, R. A.: Studies of single aerosol particles containing malonic acid, glutaric acid, and their mixtures with sodium chloride. I. Hygroscopic growth, J. Phys. Chem. A, 114(16), 5335– 5341, doi:10.1021/jp100059k, 2010.
- 30 Pratt, K. A., Hatch, L. E. and Prather, K. A.: Seasonal volatility dependence of ambient particle phase amines, Environ. Sci. Technol., 43(14), 5276–5281, doi:10.1021/es803189n, 2009. Qiu, C. and Zhang, R.: Physiochemical properties of alkylaminium sulfates: Hygroscopicity, thermostability, and density, Environ. Sci. Technol., 46(8), 4474–4480, doi:10.1021/es3004377, 2012. Qiu, C. and Zhang, R.: Multiphase chemistry of atmospheric amines., Phys. Chem. Chem. Phys. PCCP, 15(16), 5738–52,

21

doi:10.1039/c3cp43446j, 2013.

Rickards, A. M. J., Miles, R. E. H., Davies, J. F., Marshall, F. H. and Reid, J. P.: Measurements of the sensitivity of aerosol hygroscopicity and the κ parameter to the O/C ratio, J. Phys. Chem. A, 117(51), 14120–14131, doi:10.1021/jp407991n, 2013. Robinson, R. A. and Stokes, R. H.: Electrolyte Solutions, Butterworth & Co., London., 1970.

- 5 Rovelli, G., Miles, R. E. H., Reid, J. P. and Clegg, S. L.: Accurate Measurements of Aerosol Hygroscopic Growth Over a Wide Range in Relative Humidity, J. Phys. Chem. A, 120, 4376–4388, doi:10.1021/acs.jpca.6b04194, 2016. Sauerwein, M., Clegg, S. L. and Chan, C. K.: Water Activities and Osmotic Coefficients of Aqueous Solutions of Five Alkylaminium Sulfates and Their Mixtures with H2SO4 at 25oC, Aerosol Sci. Technol., 49(8), 566–579, doi:10.1080/02786826.2015.1043045, 2015.
- 10 Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, Wiley., 2006. Shiraiwa, M., Zuend, A., Bertram, A. K. and Seinfeld, J. H.: Gas-particle partitioning of atmospheric aerosols: interplay of physical state, non-ideal mixing and morphology., Phys. Chem. Chem. Phys., 15(27), 11441–53, doi:10.1039/c3cp51595h, 2013.

Shulman, M. L., Charlson, R. J. and Davis, E. J.: The effects of atmospheric organics on aqueous droplet evaporation, J. 15 Aerosol Sci., 28(5), 737–752, doi:10.1016/S0021-8502(96)00469-7, 1997.

Smith, J. N., Barsanti, K. C., Friedli, H. R., Ehn, M., Kulmala, M., Collins, D. R., Scheckman, J. H., Williams, B. J. and McMurry, P. H.: Observations of aminium salts in atmospheric nanoparticles and possible climatic implications, Proc. Natl. Acad. Sci., 107(15), 6634–6639, doi:10.1073/pnas.0912127107, 2010.

Topping, D., Connolly, P. and McFiggans, G.: Cloud droplet number enhanced by co-condensation of organic vapours, Nat. 20 Geosci., 6, 443–446, doi:10.1038/ngeo1809, 2013a.

- Topping, D., Connolly, P. and McFiggans, G.: Cloud droplet number enhanced by co-condensation of organic vapours, Nat. Geosci., 6(6), 443–446, doi:10.1038/ngeo1809, 2013b. Topping, D., Barley, M. and McFiggans, G.: Including phase separation in a unified model to calculate partitioning of vapours to mixed inorganic–organic aerosol particles, Faraday Discuss., 165, 273, doi:10.1039/c3fd00047h, 2013c.
- 25 Wang, L., Khalizov, A. F., Zheng, J., Xu, W., Ma, Y., Lal, V. and Zhang, R.: Atmospheric nanoparticles formed from heterogeneous reactions of organics, Nat. Geosci., 3(4), 238–242, doi:10.1038/ngeo778, 2010a. Wang, L., Vinita, L., Khalizov, A. F. and Zhang, R.: Heterogeneous chemistry of alkylamines with sulfuric acid: Implications for atmospheric formation of alkylaminium sulfates, Environ. Sci. Technol., 44(7), 2461–2465, doi:10.1021/es9036868, 2010b.
- 30 Wex, H., Petters, M. D., Carrico, C. M., Hallbauer, E., Massling, A., McMeeking, G. R., Poulain, L., Wu, Z., Kreidenweis, S. M. and Stratmann, F.: Towards closing the gap between hygroscopic growth and activation for secondary organic aerosol: Part 1 – Evidence from measurements, Atmos. Chem. Phys. Discuss., 9, 3987–3997, doi:10.5194/acpd-9-955-2009, 2009. Wexler, A. S. and Clegg, S. L.: Atmospheric aerosol models for systems including the ions H+, NH4+, Na+, SO42−, NO3−, Cl−, Br−, and H2O, J. Geophys. Res. Atmos., 107(D14), 4207, doi:10.1029/2001JD000451, 2002.

Yli-Juuti, T., Barsanti, K., Hildebrandt Ruiz, L., Kieloaho, a. J., Makkonen, U., Petäjä, T., Ruuskanen, T., Kulmala, M. and Riipinen, I.: Model for acid-base chemistry in nanoparticle growth (MABNAG), Atmos. Chem. Phys., 13(24), 12507–12524, doi:10.5194/acp-13-12507-2013, 2013.

You, Y., Smith, M. L., Song, M., Martin, S. T. and Bertram, A. K.: Liquid–liquid phase separation in atmospherically relevant

5 particles consisting of organic species and inorganic salts, Int. Rev. Phys. Chem., 33(1), 43–77, doi:10.1080/0144235X.2014.890786, 2014.

Zobrist, B., Soonsin, V., Luo, B. P., Krieger, U. K., Marcolli, C., Peter, T. and Koop, T.: Ultra-slow water diffusion in aqueous sucrose glasses., Phys. Chem. Chem. Phys., 13(8), 3514–3526, doi:10.1039/c0cp01273d, 2011.

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5 Table 1: Parameters of the 3rd order polynomial parameterisation for density and for the molar mixing rule (Sect. 2.1). DMAS is the **original fitting of experimental data (black circles in Fig. 9), while 'Error+' and 'Error-' are calculated supposing a ±2% error on the extrapolated ρ***melt***, as described in the text.**

Figures

Figure 1: Schematics representing the retrieval of hygroscopic growth curves from comparative kinetics experiments in a CK-EDB. 5 **An experimental sequence of alternating single evaporating probe and sample droplets is collected (a). For each pair of probe and sample droplets, the gas phase RH is inferred from the evaporation kinetics of the probe (b) and this information is used to analyse the corresponding sample droplet hygroscopic properties as indicated in (c).**

Figure 2: Measured $n_{\text{water}}/n_{\text{solute}}$ vs. a_w (Panel (a)) and osmotic coefficients (ϕ_{st}) vs. the square root of sulphate molality (m(SO4²⁻)^{0.5}) **(Panel (b)) of ammonium sulphate solution obtained from the reaction between NH³ and H2SO4. Symbols: black dots – 0.03** *mfs* **of** (NH₄)₂SO₄ in the initial solution, 88.5% RH in the gas phase; open circles - 0.004 mfs of (NH₄)₂SO₄ in the initial solution, 90% RH 5 **in the gas phase; solid lines – calculations from E-AIM model; dashed lines – uncertainty on the osmotic coefficients corresponding**

to an error in a_w of ± 0.002 .

Figure 3: Aminium sulphates GF_r vs. a_w (Panel (a)) and $n_{\text{water}}/n_{\text{solute}}$ vs. a_w (Panel (b)) hygroscopic growth plots from CK-EDB **experiments. Symbols: ● – MMAS; ■ – DMAS; ♦ – TMAS; ○ – MEAS; □ – DEAS;** ♢ **– TEAS; line, E-AIM model calculation for** (NH₄)₂SO₄. Note that only E-AIM predictions for (NH₄)₂SO₄ are reported in this and subsequent figures with the CK-EDB 5 **measurements and the E-AIM model compared in Figure 2.**

Figure 4**: Moles of salt (***nsalt***) in a 1 µm radius solid particle as a function of the molecular weight (***Mw***) of each aminium sulphate.**

Figure 5: *nwater/nsolute* **vs.** *a^w* **plots for the six aminium sulphates. Symbols: black circles – CK-EDB comparative kinetics measurements; open circles – Sauerwein et al. (2015); grey dots – Clegg et al. (2013) (based on measurements from Qiu and Zhang (2012)); line – (NH4)2SO⁴ calculation from E-AIM model.**

Figure 6: Osmotic coefficients (ϕ_{st}) vs. square root of sulphate mass fraction ($m(SO4^{2})^{0.5}$) for the six aminium sulphates. Symbols: **black circles – CK-EDB comparative kinetics measurements; open circles – Sauerwein et al. (2015); grey circles – Clegg et al. (2013) (based on measurements from Qiu and Zhang (2012)); solid line – (NH4)2SO⁴ calculation from E-AIM model; dashed lines –** 5 uncertainty of the osmotic coefficients for $(NH_4)_2SO_4$ corresponding to an error in a_w of ± 0.002 , included to provide a guide as to the **level of expected error in the osmotic coefficient with varying molality.**

Figure 7: Mass of H2O (kg) per mole of solute as a function of the degree of neutralisation of sulphuric acid by amine (*xsalt***). Symbols: solid circles – data from this work; open circles – data from Sauerwein et al. (2015). Lines: black – ZSR fitting of the data in Sauerwein et al. (2015); red dashed – same ZSR fitting but including the CK-EDB data point for each** *aw***; shaded envelopes –** 5 **uncertainty associated to the fitting.**

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Figure 8: Evaluation of the reproducibility of measurements of the hygroscopic properties of DMAS. Symbols: black dots, dark grey diamonds, light grey triangles – different datasets measured with the CK-EDB from DMAS solutions obtained indipendently; open circles – data from water activity measurements in Sauerwein et al. (2015); solid lines – (NH4)2SO⁴ calculations from the E-AIM 5 model; dashed lines – uncertainty on $(NH_4)_2SO_4$ osmotic coefficients corresponding to an error in a_w of ± 0.002 .

Figure 9: DMAS density (*ρ***) dependence on mass fraction (***mfs0.5***) of solute. Symbols: black squares – measured densities, this work; solid line – 3rd order polynomial fit of measured ρ values; grey shaded area – evaluated uncertainty of the density parameterisation (calculated as discussed in the main text); open circles – measured densities from Clegg et al. (2013); dashed line – densities from** 5 **the apparent molar volumes fitting by Clegg et al. (2013).**

Figure 10: Hygroscopic properties of DMAS obtained from the treatment of the CK-EDB kinetics evaporation data with the three different set of parameters in Table 1 for treating solution density. Symbols: black squares – original data; dark grey dots – obtained with 'Error+' parameters; light grey diamonds – obtained with 'Error-' parameters; open circles – Sauerwein et al. (2015); solid 5 lines – E-AIM model for (NH₄)₂SO₄; dashed lines – uncertainty on (NH₄)₂SO₄ osmotic coefficients corresponding to an error in a_w **of ±0.002.**

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