

1 **Impact of Intercontinental Pollution Transport on North American Ozone Air Pollution:**
2 **An HTAP Phase 2 Multi-model Study**

3
4 Min Huang^{1,2}, Gregory R. Carmichael³, R. Bradley Pierce⁴, Duseong S. Jo⁵, Rokjin J. Park⁵,
5 Johannes Flemming⁶, Louisa K. Emmons⁷, Kevin W. Bowman⁸, Daven K. Henze⁹, Yanko Davila⁹,
6 Kengo Sudo¹⁰, Jan Eiof Jonson¹¹, Marianne Tronstad Lund¹², Greet Janssens-Maenhout¹³,
7 Frank J. Dentener¹³, Terry J. Keating¹⁴, Hilke Oetjen^{8,*}, Vivienne H. Payne⁸

8
9 ¹George Mason University, Fairfax, VA, USA

10 ²University of Maryland, College Park, MD, USA

11 ³University of Iowa, Iowa City, IA, USA

12 ⁴NOAA National Environmental Satellite, Data, and Information Service, Madison, WI, USA

13 ⁵Seoul National University, Seoul, Korea

14 ⁶European Center for Medium range Weather Forecasting, Reading, UK

15 ⁷National Center for Atmospheric Research, Boulder, CO, USA

16 ⁸Jet Propulsion Laboratory, California Institute of Technology, Pasadena, CA, USA

17 ⁹University of Colorado-Boulder, Boulder, CO, USA

18 ¹⁰Nagoya University, Furocho, Chigusa-ku, Nagoya, Japan

19 ¹¹Norwegian Meteorological Institute, Oslo, Norway

20 ¹²Center for International Climate and Environmental Research, Oslo, Norway

21 ¹³European Commission, Joint Research Center, Ispra, Italy

22 ¹⁴US Environmental Protection Agency, Washington, DC, USA

23 *Now at: University of Leicester, Leicester, UK

24
25 *Correspondence to:* Min Huang (mhuang10@gmu.edu)

26 **Abstract**

27
28 The recent update on the US National Ambient Air Quality Standards of the ground-level
29 ozone (O₃) can benefit from a better understanding of its source contributions in different US
30 regions during recent years. In the Hemispheric Transport of Air Pollution experiment Phase 1
31 (HTAP1), various global models were used to determine the O₃ source-receptor relationships
32 among three continents in the Northern Hemisphere in 2001. In support of the HTAP Phase 2
33 (HTAP2) experiment that studies more recent years and involves higher-resolution global models
34 and regional models' participation, we conduct a number of regional scale Sulfur Transport and
35 dEposition Model (STEM) air quality base and sensitivity simulations over North America during
36 May-June 2010. STEM's top and lateral chemical boundary conditions were downscaled from
37 three global chemical transport models' (i.e., GEOS-Chem, RAQMS, and ECMWF C-IFS) base
38 and sensitivity simulations in which the East Asian (EAS) anthropogenic emissions were reduced
39 by 20%. The mean differences between STEM surface O₃ sensitivities to the emission changes
40 and its corresponding boundary condition model's are smaller than those among its boundary
41 condition models, in terms of the regional/period mean (<10%) and the spatial distributions. An
42 additional STEM simulation was performed in which the boundary conditions were downscaled
43 from a RAQMS simulation without EAS anthropogenic emissions. The scalability of O₃
44 sensitivities to the size of the emission perturbation is spatially varying, and the full source
45 contribution obtained by linearly scaling the North American mean O₃ sensitivities to a 20%
46 reduction in the EAS anthropogenic emissions may be underestimated by at least 10%.

47 The three boundary condition models' mean O₃ sensitivities are ~8% (May-June 2010)/~11%
48 (2010 annual) lower than those estimated by multiple global models, and the multi-model
49 ensemble estimates are higher than the HTAP1 reported 2001 conditions, due to the growing EAS
50 anthropogenic emissions, the interannual variability in atmospheric circulation (i.e., stronger trans-
51 Pacific transport in spring 2010 following an El Niño event), and the different experiment designs
52 of HTAP1 and HTAP2. GEOS-Chem sensitivities indicate that the EAS anthropogenic NO_x
53 emissions matter more than the other EAS O₃ precursors to the North American O₃, qualitatively
54 consistent with previous adjoint sensitivity calculations.

55 In addition to the analyses on large spatial/temporal scales relative to the HTAP1, we also
56 show results on subcontinental- and event-scale that are more relevant to the US air quality
57 management. Satellite O₃ (TES, JPL-IASI, and AIRS) and carbon monoxide (TES and AIRS)
58 products, along with surface measurements and model calculations, show that during certain
59 episodes stratospheric O₃ intrusions and the transported EAS pollution influenced O₃ in the western
60 and the eastern US differently. Free-running global models underpredicted the transported
61 background O₃ during these episodes, posing difficulties for STEM to accurately simulate the
62 surface O₃ and its source contribution. Although we effectively improved the modeled O₃ by
63 incorporating satellite O₃ (OMI and MLS) and evaluated the quality of the HTAP2 emission
64 inventory with the KNMI OMI nitrogen dioxide, using observations to evaluate and improve O₃
65 source attribution still remains to be further explored.

1. Introduction

Tropospheric ozone (O_3), a short-lived trace gas with a lifetime ranging from hours in the boundary layer to weeks in the free troposphere, affects tropospheric chemistry, harms human and ecosystem health, and induces climate change on local, regional and global scales (Jerrett et al., 2009; Smith et al., 2009; Anenberg et al., 2010; Mauzerall and Wang, 2001; Avnery et al., 2011a, b; Shindell et al., 2009, 2013; Bowman and Henze, 2012; Stevenson et al., 2006, 2013; Monks et al., 2015). It has been recognized that the uneven distributions of tropospheric O_3 can be attributed to the stratosphere as well as local, regional and distant emission sources, through complicated processes that occur on synoptic, meso- and micro-scales (Task Force on Hemispheric Transport of Air Pollution (HTAP), 2010; National Research Council (NRC), 2009; Maas and Grennfelt, 2016). The mitigation of O_3 's climate and health impacts would benefit from efforts to control the emissions of its precursors from the various emission sources (United Nations Environment Programme (UNEP) and World Meteorological Organization (WMO), 2011), such as nitrogen oxides (NO_x), carbon monoxide (CO), methane (CH_4), and non-methane volatile organic compounds (NMVOCs).

Ground-level O_3 is one of the six criteria air pollutants regulated by the US Environmental Protection Agency (EPA), and the US National Ambient Air Quality Standards (NAAQS) has recently been lowered to 70 ppbv to better protect Americans' health and the environment. Issues regarding making accurate estimates of the total O_3 as well as the background O_3 level (defined as the concentration that is not affected by recent locally-emitted or produced anthropogenic pollution) (e.g., McDonald-Buller et al., 2011; Zhang et al., 2011; Fiore et al., 2014; Huang et al., 2015), have been recently discussed as part of the implementation of the new US O_3 standard (US EPA, 2016a, b). This includes assessing the impacts of various components of the background O_3 , such as stratospheric O_3 , local natural sources such as biogenic, lightning and wildfire emissions, as well as the long-range transport (LRT) of pollution. The impact of the trans-Pacific pollution transport on US air quality has been evaluated in numerous studies over the past decades (e.g., Fiore et al., 2009; Reidmiller et al., 2009; Zhang et al., 2008, 2009; Huang et al., 2010, 2013a; Lin et al., 2012a, 2015, 2016; US EPA, 2016a). It has been found that the increasing trends of pollution in the upwind continents, especially the populated East Asia (e.g., Zhang et al., 2014; Susaya et al., 2013; Wang et al., 2012), may partially offset the US air quality improvements in recent decades due to the regional and local emission controls (e.g., Jacob et al., 1999; Verstraeten et al., 2015; Ambrose et al., 2011; Wigder et al., 2013; Cooper et al., 2010; Parrish et al., 2009, 2012; Gratz et al., 2014). A better understanding of the processes that determine the O_3 pollution levels, as well as an improved capability of attributing the air pollution to nearby or distant sources is needed to assist with designing and implementing effective local emission control strategies to comply with the tighter air quality standards.

Chemical transport models are often used to reproduce and attribute the observed O_3 levels, including assessing the impacts of the internationally transported O_3 on the US air quality. In the HTAP modeling experiment Phase 1 (HTAP1), various global models with horizontal resolutions ranging from $1^\circ \times 1^\circ$ to $5^\circ \times 5^\circ$, only around half of which are finer than $3^\circ \times 3^\circ$, were used to determine the O_3 source-receptor (SR) relationships among three continents in the Northern Hemisphere in 2001 (Chapter 4 in HTAP, 2010). The global model based SR relationships in HTAP1 determined using the emission perturbation approach (i.e., calculating the changes of O_3

112 at the receptor regions in response to a 20% reduction in the emission inputs in a given source
113 region) were reported as either monthly 24h mean values or policy-relevant metrics such as the
114 maximum daily 8h average (MDA8) for the US (e.g., Fiore et al., 2009; Reidmiller et al., 2009).
115 Large intermodel diversity was found in the simulated total O₃ and the intercontinentally
116 transported pollution for the chosen SR pairs in the northern midlatitudes, indicating the challenges
117 with model simulations to accurately represent the key atmospheric processes. Multi-model mean
118 results were the foci of in these studies with the assumption that this approach can reduce the
119 uncertainty from the single model estimates for monthly or seasonal means. “Ensemble” model
120 analyses have been suggested by some US stakeholders as one of the methods for helping with the
121 characterization of the background O₃ components (US EPA, 2016b). Although the multi-model
122 approach can help identify some of the weaknesses of the individual models and may produce
123 more reliable estimates, it is necessary to well understand the uncertainties inherent in using the
124 same set of anthropogenic emissions in all these model simulations. Satellite observations over the
125 regions with limited in-situ measurements such as the East Asia can be particularly helpful for
126 quantifying such uncertainties.

127
128 The 20% emission perturbation in the HTAP1 modeling experiment was chosen to produce
129 a sizeable (i.e., larger than numerical noise) and realistic impact, but small enough in the assumed
130 near-linear atmospheric chemistry regime. The scalability of the modeled O₃ sensitivities to the
131 size of the emission perturbations has been assessed on continental scale (Wu et al., 2009; Fiore et
132 al., 2009; HTAP, 2010; Wild et al., 2012; Emmons et al., 2012). The receptor O₃ responses to the
133 source-region emission perturbations are found to be fairly linear within ~50% of the perturbations.
134 However, due to the chemical non-linearity, the full source contribution obtained by linearly
135 scaling the receptor regional mean O₃ sensitivity to the 20% reduction in the source region
136 emissions may be underestimated, and the scalability depended on seasons and the perturbed
137 emission species. Huang et al. (2013b) investigated the scalability of the O₃ sensitivity between
138 the southern California-US intermountain west SR pair for May 2010, in which study the southern
139 California anthropogenic emissions were perturbed by multiple amounts of +50%, -50%, -100%.
140 They reported that the scalability of the O₃ sensitivities changed with the distance from the source
141 regions. Further analyses on the scalability of these modeled O₃ sensitivities during recent years
142 especially for the East Asia-NAM SR pair, as well as their spatial variability, are still needed.
143 Furthermore, results generated using the emission perturbation approach need to be compared with
144 those based on the other methods (e.g., tagged tracers, adjoint sensitivity).

145
146 Previous studies have demonstrated the advantages of high resolution chemical transport
147 modeling for understanding SR relationships (e.g., Lin et al., 2010 for Europe and the East Asia;
148 Lin et al., 2012a; Huang et al., 2010, 2013a for Asia and NAM). Using observations (satellite,
149 sondes, aircraft) along with single model simulations, a few studies have reported that the US O₃
150 sensitivities to extra-regional sources is time- and region-dependent (e.g., Lin et al., 2012a, b;
151 Langford et al., 2011; Ott et al., 2016), and therefore the necessity of evaluating the extra-regional
152 source impacts on event scale has been emphasized in these studies as well as in US EPA (2016a,
153 b). The HTAP Phase 2 (HTAP2) multi-model experiment, initiated in 2012, is designed to advance
154 the understanding of the impact of intercontinental pollution transport during more recent years
155 (i.e., 2008-2010) involving a number of global and regional models’ participation (Galmarini et
156 al., 2017; Koffi et al., 2016). The regional models are anticipated to help connect the analyses over
157 global and regional scales and enable discussions on small spatial (e.g., subcontinental) and

158 temporal scales (i.e., event based analyses). The use of satellite products for identifying the
159 transport events as well as for quantitative model evaluation is also encouraged in the work plan.
160 The HTAP2 modeling experiment was sequentially conducted in two steps. First, similar to the
161 HTAP1 experiment, a group of global models with different resolutions conducted base and
162 emission perturbation sensitivity simulations to determine the pollutants' SR relationships. All
163 models in their base simulations used the same set of harmonized sector-based global
164 anthropogenic emissions developed specifically for the HTAP2 modeling experiment (Janssens-
165 Maenhout et al., 2015). Most of these global models recorded only key chemical species from their
166 base and sensitivity simulations in varied temporal frequencies. Several global models saved the
167 three-dimensional (3D) chemical fields of more species with a 3- or 6-hour interval, which are
168 suitable for being used as regional models' chemical boundary conditions. In the second step,
169 regional models conducted base and sensitivity simulations to analyze the pollutants' SR
170 relationships in greater detail. The regional model simulations used the same set of anthropogenic
171 emissions as the global models within their simulation domains, and the chemical boundary
172 conditions in these regional simulations were downscaled from the base and sensitivity simulations
173 from the selected boundary condition model outputs. For regional simulations over the North
174 America and Europe, boundary conditions were mostly taken from a single model such as the
175 ECMWF C-IFS or GEOS-Chem.

176
177 This study aims to address: 1) the differences in O₃ sensitivities generated from the HTAP2
178 and HTAP1 experiments to help address how the LRT impacts on NAM changed through time; 2)
179 how the refined modeling experiment design in HTAP2 can help advance our understanding of the
180 LRT impacts on NAM, particularly the involvement of regional models and the inclusion of small
181 spatial/temporal scale analysis during high O₃ episodes that are more relevant to air quality
182 management; 3) the usefulness of satellite observations for better understanding the sources of
183 uncertainties in the modeled total O₃ (e.g., from the emission and regional models' boundary
184 condition inputs) as well as for reducing the uncertainties in some of these model inputs via
185 chemical data assimilation. We performed a number of regional scale STEM (Sulfur Transport
186 and dEposition Model) base and sensitivity simulations over the NAM during May-June 2010,
187 during which period strong trans-Pacific pollution transport were shown to episodically impact the
188 US (Lin et al., 2012a). Extending the HTAP2 regional simulations' basic setup, the STEM top and
189 lateral chemical boundary conditions were downscaled from three global models' (i.e., the Seoul
190 National University (SNU) GEOS-Chem, RAQMS, and the ECMWF C-IFS) base and sensitivity
191 simulations in which the East Asian anthropogenic emissions were reduced. The STEM surface
192 O₃ sensitivities over the NAM region based on different boundary condition models were inter-
193 compared, in terms of the regional averages and the spatial patterns on monthly basis and during
194 a selected event identified by satellite O₃ and CO products. These were also compared with the
195 sensitivities estimated by their corresponding boundary condition models as well as all HTAP2
196 participating global models and the results from HTAP1.

197 198 **2. Methods**

199 *2.1. Anthropogenic emission inputs*

200
201 Identical anthropogenic emissions were used in all global and regional chemical transport
202 models' base and sensitivity simulations. This monthly-varying harmonized sectoral (i.e., power,
203 industry, transportation, residential, shipping, aircraft, agriculture) emission inventory was

204 provided on a gridded $0.1^\circ \times 0.1^\circ$ resolution for the years of 2008 and 2010, by compiling the
205 officially reported emissions at the national scale (Janssens-Maenhout et al., 2015;
206 http://edgar.jrc.ec.europa.eu/htap_v2). The temporal profiles for developing the monthly-varying
207 emissions differ by region and sector. The amount of emissions of key O₃ precursors (CO, NO_x,
208 NMVOCs) from both years are summarized in Table S1 for the four major emissions sectors, over
209 the NAM (US+Canada, based on data from the US EPA and the Environmental Canada, which
210 shows lower emissions from the previous years as also discussed in Pouliot et al., 2015), MICS-
211 Asia regions (south, southeast, and east Asia, based on country inventory for China and from the
212 Clean Air Policy Support System and the Regional Emission inventory in ASia 2.1, more
213 information also in Li et al., 2017), and for over the world. For all of these species, global total
214 emissions in 2008 and 2010 are similar. The NO_x, NMVOC, and CO emissions decreased from
215 2008 to 2010 over the NAM by 10.7%, 9.4%, and 15.7%, respectively. In 2008, NAM NO_x,
216 NMVOC and CO contributed to 18.0%, 11.7% and 11.9% of the global total, respectively, and in
217 2010, these contributions became 15.8%, 10.5% and 10.2%. For 2010, the transportation sector
218 contributed more than the other sectors to NAM anthropogenic NO_x and CO emissions; industrial
219 sector contributed more than the other sectors to NMVOCs emissions. Over East Asian countries,
220 these emissions are ~2-5 times higher than the US emissions, and the NO_x, NMVOC and CO
221 emissions increased over Asia by 7.3%, 7.2% and 1.0%, with the dominant emission sectors in
222 2010 of transportation, industry, and residential, respectively. For both years, the emissions over
223 the MICS-Asia regions contribute to over 40% of the global emissions. For these key O₃ precursors,
224 the East Asian countries contribute to 45% (NMVOCs)-70% (NO_x) of the emissions in the MICS-
225 Asia domain in both years, and the south Asian countries contribute to ~22% (NO_x)-34%
226 (NMVOCs) of the MICS-Asia emissions. The uncertainty of the emission estimates differs by
227 emission sector and species: i.e., the emissions from large-scale combustion sources (e.g., NO_x
228 and CO from power and industry sectors) are less uncertain than those from small-scale and
229 scattered sources (e.g., CO and NMVOCs from transportation and residential sources). Non-
230 anthropogenic emission inputs used in different models' simulations may differ, and their impacts
231 on the modeled total O₃ and the SR relationships will be compared in detail in future studies.

232

233 2.2. *Region definitions for the SR study and the model base and sensitivity simulations*

234 2.2.1. Base and 20% emission perturbation simulations from global and regional models

235 The HTAP2 simulations from eight global models, used in this study, are listed in Table
236 1a, including the relevant references. Horizontal and vertical resolutions of these models range
237 from finer than 1° to coarser than 2.5° , and from 20 to 60 layers, respectively. Overall these
238 resolutions are higher than the HTAP1 participating models'. Figure 1 defines the source regions
239 used in the HTAP2 SR relationship study and we will focus in this study on assessing the East
240 Asia (EAS), S Asia (SAS), Europe (EUR), and non-NAM anthropogenic source (interchangeable
241 in this paper with "(all) foreign") impacts on the NAM O₃ levels in 2010. Specifically, each model
242 performed a base simulation and a number of sensitivity simulations in which the original HTAP2
243 anthropogenic emissions for all species and sectors in a defined source region were perturbed by
244 a certain amount (referring to 20% as in most cases) and these cases are defined in Table 1a-b as
245 **source region*ALL(*perturbation*)*, where "ALL" refers to "all species and sectors", consistent
246 with HTAP1 and HTAP2's naming convention. The O₃ differences $R(O_3, *source region*,$
247 **perturbation*)* over the NAM were then calculated between each model's base and sensitivity
248 simulations:

249

$$\begin{aligned}
250 \quad R(\text{O}_3, \text{EAS}, 20\%) &= \text{BASE O}_3 - \text{EASALL}(-20\%) \text{ O}_3 & (1a) \\
251 \quad R(\text{O}_3, \text{SAS}, 20\%) &= \text{BASE O}_3 - \text{SASALL}(-20\%) \text{ O}_3 & (1b) \\
252 \quad R(\text{O}_3, \text{EUR}, 20\%) &= \text{BASE O}_3 - \text{EURALL}(-20\%) \text{ O}_3 & (1c) \\
253 \quad R(\text{O}_3, \text{non-NAM}, 20\%) &= \text{NAMALL O}_3 - \text{GLOALL}(-20\%) \text{ O}_3 & (1d)
\end{aligned}$$

254
255 The monthly-mean $R(\text{O}_3, \textit{*source region*}, 20\%)$ values were averaged over the NAM
256 region for the analysis and compared with the findings in the HTAP1 study (e.g., Fiore et al., 2009).
257 It is worth mentioning that the rectangular source regions defined in HTAP1 were modified in
258 HTAP2 to align with the geo-political borders. For EAS and SAS, the regions not overlapped by
259 HTAP1 and HTAP2 are mostly in the less populated/polluted regions such as the northwestern
260 China, according to the HTAP2 emission maps (http://edgar.jrc.ec.europa.eu/htap_v2/index.php).
261 HTAP2's EUR domain excludes certain regions in Russia/Belarussia/Ukraine, Middle East and
262 North Africa that are included in HTAP1's EUR domain. The impact of emissions over these
263 regions on comparing the NAM $R(\text{O}_3, \text{EUR}, 20\%)$ values in HTAP1 and HTAP2 will be discussed
264 in Section 3.2.1.

265
266 A unitless “Response to Extra-Regional Emission Reductions (RERER)” metric
267 (Galmarini et al., 2017), as defined in eq. (2), was also calculated to measure the importance of
268 local versus non-local sources to NAM's O_3 levels:

$$269 \quad \text{RERER}(\text{O}_3, \text{NAM}) = \frac{R_{\text{O}_3, \text{non-NAM}, 20\%}}{R_{\text{O}_3, \text{global}, 20\%}} = \frac{(\text{NAMALL O}_3 - \text{GLOALL O}_3)}{(\text{BASE O}_3 - \text{GLOALL O}_3)} \quad (2)$$

270 The denominator and numerator terms of RERER represent the impacts of global and non-NAM
271 anthropogenic emissions on NAM O_3 , respectively. The higher the NAM RERER value is, the
272 stronger impact from non-local sources on NAM is indicated. The RERER value can exceed 1,
273 when emission reductions led to increasing concentrations (e.g. O_3 titration by nitrogen monoxide
274 (NO)).

275
276 The STEM (version 2K3) regional simulations were then performed on a 60 km×60 km
277 horizontal resolution (a typical coarse regional model resolution) grid over NAM within the
278 domain defined in Figure 2a during May-June 2010. The meteorological conditions in spring 2010
279 were compared with the climatology from the NCEP/NCAR reanalysis data for the 1981-2010
280 period (Kalnay et al., 1996) in Huang et al. (2013b), concluding that this spring represents a period
281 of stronger-than-climatological average spring trans-Pacific transport, based on a stronger
282 meridional gradient in the North Pacific and higher Pacific/North American (PNA) indexes. This
283 is consistent with the findings by Lin et al. (2014) that the El Niño conditions during the 09/10
284 winter strengthened the trans-Pacific transport of Asian pollution in spring 2010. The mean near-
285 surface air temperatures in the western US in this spring were lower than the climatology, with
286 larger anomalies in the mountain states, which may have led to weaker local O_3 production and
287 decomposition of the transported peroxyacyl nitrates (PAN). In contrast, higher-than-normal
288 temperatures were found in the eastern US that favored anomalously strong local O_3 production.

289
290 STEM has been used to interpret the observations collected by satellites and during aircraft
291 campaigns in the past decade (e.g., Carmichael et al., 2003a, b; Huang et al., 2010, 2013a, b, 2014,
292 2015). STEM calculates gas-phase chemistry reactions based on the SAPRC 99 gaseous chemical
293 mechanism (Carter, 2000) with thirty photolysis rates calculated online by the Tropospheric
294 Ultraviolet-Visible radiation model (Madronich et al., 2002). Most of the key configurations of the
295 60 km base simulations are the same as those described in Lapina et al. (2014), i.e., meteorological

296 fields were pre-calculated by the Advanced Research Weather Research and Forecasting Model
297 (WRF-ARW, Skamarock et al., 2008) version 3.3.1 forced by the North American Regional
298 Reanalysis data (Mesinger et al., 2006), using a similar set of the physics configuration to those in
299 Huang et al. (2013a). Biomass burning emissions are from the Fire INventory from NCAR (FINN)
300 inventory version 1.0 (Wiedinmyer et al., 2011). Biogenic emissions were calculated by the Model
301 of Emissions of Gases and Aerosols from Nature (MEGAN) version 2.1 (Guenther et al., 2012),
302 driven by the WRF meteorology. Lightning NO_x emissions are generated following the method in
303 Allen et al. (2012), with the flash rates determined by the WRF convective precipitation and scaled
304 to the National Lightning Detection Network flash rates. A major difference of the STEM
305 simulations in this study from the Lapina (2014) study is that the anthropogenic emissions were
306 replaced with the monthly-mean HTAP2 inventory with no weekday-weekend variability applied,
307 rather than the earlier National Emission Inventory (NEI) 2005 in which the weekday-weekend
308 variability exists. This change can introduce uncertainty for some US regions where weekday-
309 weekend variability of some O₃ precursors' emissions was notable during the studied period (e.g.,
310 weekend NO_x emissions in southern California during spring/summer 2010 were 0.6-0.7 of the
311 weekday emissions as reported by Kim et al. (2016) and Brioude et al. (2013)), but this was done
312 to ensure consistency with the HTAP2 global model simulations, that also didn't use daily variable
313 emissions for any regions in the world. The VOC speciation for the SPRAC 99 chemical
314 mechanism in the NEI 2005 (ftp://aftp.fsl.noaa.gov/divisions/taq/emissions_data_2005) were
315 applied to break down the total NMVOC emissions provided in the HTAP2 inventory. The VOC
316 speciation based on the year of 2005 can be unrealistic for 2005 as well as 2010 as studies have
317 reported variable temporal changes of different VOC species in some US cities (e.g., Warneke et
318 al., 2012). The time-varying lateral and top boundary conditions in the STEM base simulations
319 were downscaled from three global models (i.e., 3 hourly SNU GEOS-Chem, 3 hourly ECMWF
320 C-IFS, and 6 hourly RAQMS) base simulations. In support of the SR relationship study to quantify
321 the East Asia anthropogenic impacts on the NAM, three STEM sensitivity simulations were also
322 conducted in which the STEM boundary conditions were downscaled from the EASALL(-20%)
323 sensitivity simulations by these three global models (Table 1b). All STEM simulated 3D chemical
324 fields were saved hourly for the convenience of calculating the US primary O₃ standard metric
325 MDA8 as well as the quantitative comparisons against the satellite Level 2 (L2) O₃ products. The
326 STEM base case surface O₃ performance and its O₃ sensitivities were also compared with those of
327 its boundary condition models as well as the multi- global model means. The latitude/longitude
328 ranges (20-50°N/130-65°W) of NAM for the global and regional model based sensitivity
329 calculations were selected to mainly account for the coverage of the STEM domain, which are
330 slightly different from the definition of North America in HTAP1.

331
332 Note that non-anthropogenic emission inputs used in STEM and its boundary condition
333 models differed, as summarized in Table 1c. Figure S1 shows detailed comparisons between
334 STEM and GEOS-Chem's non-anthropogenic (i.e., soil, lightning, biomass burning) NO_x
335 emission inputs, and their impacts on the modeled NAM background O₃ were included in Lapina
336 et al. (2014). Such quantitative comparisons will also be carried out between STEM and its other
337 boundary condition models in future studies.

338
339 2.2.2. Additional base and sensitivity simulations from selected models

340

341 In addition to the base and 20% EAS all-category emission perturbation simulations, the
342 global RAQMS model conducted a sensitivity simulation in which the East Asian anthropogenic
343 emissions were zeroed out, which was also used as STEM’s boundary conditions (Table 1b). We
344 calculate the “S<sub>O₃” metric (eq. (3)) using the O₃ sensitivities in STEM and RAQMS at the receptor
345 regions in response to both 20% and 100% of emission reductions, to explore the relationships
346 between the O₃ sensitivity and the size of the emission perturbation. A closer-to-one “S<sub>O₃” value
347 indicates higher scalability of the sensitivity based on the 20% emission perturbation method for
348 obtaining the full “contribution” of the East Asian anthropogenic emissions on the NAM O₃.</sub></sub>

$$350 S_{O_3} = R(O_3, \text{EAS}, 100\%) / R(O_3, \text{EAS}, 20\%) / 5 \quad (3)$$

351 Where: $R(O_3, \text{EAS}, 100\%) = \text{BASE } O_3 - \text{EASALL}(-100\%) O_3$

352
353 The RAQMS model also provided a base simulation that assimilated satellite O₃ products
354 from the Ozone Monitoring Instrument (OMI, Levelt et al., 2006) and Microwave Limb Sounder
355 (MLS, Livesey et al., 2008) (Pierce et al., 2007), which was used to help better understand the
356 regional model base run error sources, as well as for demonstrating the use of satellite observations
357 to help improve the representation of the trans-boundary pollution.

358
359 We also used a number of sensitivity simulations produced by the GEOS-Chem adjoint
360 model v35f in which the emissions from selected anthropogenic emission sectors (power&industry,
361 transportation, residential) or individual O₃ precursor chemical species (NO_x, VOC, CO) over the
362 East Asia were reduced by 20%. Additional simulations for the 2008-2009 periods by the SNU
363 GEOS-Chem were also utilized to quantify the East Asia and non-NAM anthropogenic source
364 impacts in comparison with the 2010 conditions that we mainly focus on in this study.

365 2.3. *In-situ and satellite observations*

366 2.3.1. In-situ observations

367 Over the receptor NAM, the hourly O₃ observations at the Clean Air Status and Trends
368 Network (CASTNET, <http://epa.gov/castnet/javaweb/index.html>) sites were used to evaluate the
369 global and regional models’ base simulations in four subregions: western US (i.e., the EPA regions
370 8, 9, 10); southern US (i.e., the EPA regions 4 and 6), the Midwest (i.e., the EPA regions 5 and 7),
371 and the northeast (i.e., the EPA regions 1-3). The numbers of sites used in global and regional
372 models’ evaluation in each US subregion are summarized in Tables 2-3. The locations of these
373 sites and the subregions they belong to are indicated in Figure 2a, overlaid on a model-based terrain
374 height map. A majority of the CASTNET sites in the western US are located at high elevation (>1
375 km) remote or rural regions, more susceptible to the trans-boundary pollution (e.g., Jaffe, 2011).
376 Most of the sites in the other three subregions are located in low elevation regions, mainly affected
377 by local and regional pollution. The model-based terrain heights fairly well represent the reality
378 on subregional scale – the differences between the actual and model-based subregional mean
379 terrain heights at the CASTNET sites are smaller than 0.1 km (Table 3).

380
381 During May-June 2010, intense ozonesonde measurements were made at multiple
382 California locations (Cooper et al., 2011), in support of the NOAA “California Nexus (CalNex):
383 Research at the Nexus of Air Quality and Climate Change” field experiment (Ryerson et al., 2013).
384 They have been used to evaluate the simulated O₃ vertical profiles by the HTAP2 participating
385 models. The detailed evaluation results have been shown by Cooper et al. (2016), and will be
386 covered by subsequent publications.

387 Over HTAP2's EAS source region, the global models' O₃ performance was evaluated
388 against the monthly-mean surface in-situ O₃ measurements at 11 sites within the Acid Deposition
389 Monitoring Network in East Asia (EANET, <http://www.eanet.asia>) that had data throughout the
390 year of 2010. These include eight Japanese and three Korean sites (Figure 3a), all of which are
391 located at low elevation regions (2-150 m). The reported monthly mean observations at these sites
392 were based on weekly or daily sampled data, varying among sites.

393
394 2.3.2. Satellite products

396 In two case studies of high O₃ episodes, L2 and L3 O₃ and CO retrievals from several
397 satellite instruments were used to assess the impacts of trans-Pacific pollution transport and
398 stratospheric O₃ intrusions on NAM O₃ levels in early May. These include: 1) the early afternoon
399 O₃ and CO profiles version 5 from the Tropospheric Emission Spectrometer (TES) (Beer et al.,
400 2001; Beer, 2006) on the Aura satellite; 2) the mid-morning O₃ profiles from the METOP-Infrared
401 Atmospheric Sounding Interferometer (IASI), which were retrieved using the Jet Propulsion
402 Laboratory (JPL) TES optimal estimation retrieval algorithm (Bowman et al., 2006) for selected
403 areas including the western US (Oetjen et al., 2014, 2016); as well as 3) the early afternoon L3 O₃
404 and CO maps (version 6, 1°×1°) from the Aqua Atmospheric Infrared Sounder (AIRS) instrument.
405 The TES tropospheric O₃ retrieval is often sensitive to the mid- to lower free troposphere, and O₃
406 at these altitudes in the Eastern Pacific is known to possibly impact the downwind US surface air
407 quality at later times (Huang et al., 2010; Parrish et al., 2010). TES O₃ is generally positively
408 biased by <15% relative to high accuracy/precision reference datasets (e.g., Verstraeten et al.,
409 2013). Although IASI is in general less sensitive than TES due to its coarse spectral resolution, the
410 681–316 hPa partial column-averaged O₃ mixing ratios in the JPL product agree well with TES
411 O₃ for the 2008–2011 period with a -3.9 ppbv offset (Oetjen et al., 2016). Note that IASI O₃ data
412 are processed operationally in Europe using a different algorithm. For this work we used O₃
413 profiles from TES and IASI processed using a consistent algorithm at JPL, although the latter set
414 of data represents only a small subset of the full set of the IASI radiance measurements. The IASI
415 and TES L2 O₃ profiles (screened by the retrieval quality and the C-Curve flags) were used to
416 evaluate the STEM O₃ vertical distributions in the different base simulations, and the satellite
417 observation operators were applied in these comparisons. Taking TES as an example, its
418 observation operator h_z for O₃ is written in (4):

419
$$h_z = z_c + A_{\text{TES}} (\ln(F_{\text{TES}}(c)) - z_c) \quad (4)$$

420 where z_c is the natural log form of the TES constraint vector (a priori) in volume mixing ratio.
421 A_{TES} is the averaging kernel matrix reflecting the sensitivity of retrieval to changes in the true state
422 (Rodgers, 2000). F_{TES} projects the modeled O₃ concentration fields c to the TES grid using spatial
423 and temporal interpolation. The exponential of h_z is then used to compute the mismatches between
424 the model and TES O₃ retrievals as the model evaluation. A small mismatch between model with
425 the satellite observation operators and the satellite retrievals may indicate either good model
426 performance or may be the low sensitivity of the retrievals to the true O₃ profile. AIRS O₃ is
427 sensitive to the altitudes near the tropopause, with positive biases over the ozonesondes in the
428 upper troposphere (e.g., Bian et al., 2007); AIRS CO is most sensitive to 300–600 hPa (Warner et
429 al., 2007) and is frequently used together with the AIRS O₃ to distinguish the stratospheric O₃
430 intrusions from long-range transported anthropogenic or biomass burning pollution. We use the
431 L3 AIRS products in this study to get a broad overview of the areas that are strongly impacted by
432 the stratospheric O₃ intrusions or/and LRT of pollution.

433 The bottom-up NO_x emissions from the HTAP2 inventory were assessed on a monthly base
434 by comparing the GEOS-Chem nitrogen dioxide (NO₂) columns with the de-striped KNMI (Royal
435 Netherlands Meteorological Institute) OMI column NO₂ product version 2.0 (Boersma et al.,
436 2011a, b). For this model evaluation against the OMI L2 products, the NO₂ fields calculated by the
437 GEOS-Chem adjoint model were saved daily at 13:30 local solar time, roughly coinciding with
438 the Aura and Aqua overpassing times. Other parameters used in the model column calculations
439 came from the GEOS-5/GEOS-Chem monthly mean conditions. The OMI data that passed the
440 tropospheric quality flag at 13-14 local time were selected based on the following screening criteria:
441 surface albedo<0.3; cloud fraction<0.2; solar zenith angle <75°; and viewing zenith angle <45°.
442 The averaging kernels (Eskes and Boersma, 2003) and Air Mass Factors (AMFs) in the KNMI
443 product were used to calculate the modeled tropospheric NO₂ vertical columns comparable to the
444 OMI's. Details of the method to compare the model-based NO₂ columns with the KNMI OMI's
445 can be found in Huang et al. (2014).

446

447 **3. Results and Discussions**

448 *3.1. Evaluation of the HTAP2 bottom-up NO_x emissions and the model base simulations*

449 3.1.1. Evaluation of the bottom-up NO_x emissions

450

451 The comparison of the GEOS-Chem adjoint NO₂ columns with the OMI product was used
452 to help assess the bottom-up HTAP2 NO_x emissions. Figure 4 shows that NO₂ columns from
453 GEOS-Chem's base simulations over the US are overall overestimated. While grid-scale
454 differences in NO₂ columns may not be directly indicative of emissions biases (Qu et al., 2016),
455 these discrepancies are possibly due to a positive bias in the bottom-up emissions, mainly from the
456 anthropogenic sources, which have also been pointed out by Anderson et al. (2014) and Travis et
457 al. (2016). Larger OMI-model disagreement was found over the central/eastern US in June 2010
458 than in May, likely also due to the uncertainty in GEOS-Chem's soil or lightning NO_x emissions,
459 which appear to be high over these regions (Figure S1). The NO₂ columns in the GEOS-Chem
460 base simulation were overestimated in many northern China rural areas and underpredicted in a
461 few urban areas in the East Asia as well as a broad area in the southwestern China. The mismatches
462 between model and OMI NO₂ fell within the ranges of the comparison between the GOME2 NO₂
463 column product and six models' simulations over China in summer 2008 (Quennehen et al., 2016).
464 Also, the use of monthly-mean anthropogenic emissions as well as the overall rough treatment of
465 emission height and temporal profiles can be sources of uncertainty. These global model
466 evaluation results suggest that the EAS-NAM SR relationships analyzed using this inventory may
467 overall overestimate the NAM local contribution and underestimate the EAS contribution—Under
468 different chemical regimes, this statement would also rely on the quality of other O₃ precursors'
469 emissions in the HTAP2 inventory, and they may be associated with variable uncertainties
470 depending on the species or emission sector as introduced in Section 2.1. Therefore, careful
471 assessment of other key O₃ precursors' emissions in the inventory is needed in the future work. It
472 is important to note that uncertainty in satellite retrievals can prevent us from producing accurate
473 assessment on emissions (e.g., van Noije et al., 2006), and this comparison does not account for
474 the biases in the used OMI data, and would be further validated by using other OMI NO₂ products
475 as well as the bias-corrected (if applicable) in-situ NO₂ measurements. We also recommend more
476 global models to save their calculations more frequently, at least near the satellite overpassing
477 times, for a more comprehensive assessment of the emission inventory and a better understanding
478 of the model biases.

479 3.1.2. Evaluation of the global model O₃ performance in NAM and EAS

480

481 The monthly-mean surface O₃ from multiple global models' free runs was evaluated with
482 the CASTNET observations, at the stations with 95% of the hourly O₃ observation completeness
483 for the 1 May-30 June 2010 period. The mean biases and RMSEs for these two months were
484 summarized in Table 2a by US subregions. The three boundary condition-model as well as the
485 eight-model ensembles overall underpredicted O₃ in the western US (by ~3-6 ppbv), similar to the
486 HTAP1 model performance over these regions for May-June 2001 presented in Fiore et al. (2009).
487 This can be due to the underestimated trans-boundary pollution (as indicated by the evaluation of
488 modeled O₃ profiles with ozonesondes and satellite O₃ products). In addition, the coarser model
489 resolutions are less capable of resolving the local features that influence the pollutants' import
490 processes, chemical transformation, as well as regional processes such as the cross-state pollution
491 transport over complex terrains. The global RAQMS base simulation with satellite assimilation
492 improved the free tropospheric O₃ structure as its comparisons with the ozonesondes shows, which
493 also enhanced the simulated monthly-mean surface O₃ by up to >10 ppbv in the western US and
494 some coastal areas in the southeastern US (Figure S2, left). The global models overall significantly
495 overestimated O₃ in the other three subregions (by 8-12 ppbv), close to HTAP1 model performance
496 for May-June 2001 over the similar areas (Fiore et al., 2009) and in the Lapina et al. (2014) study
497 for 2010, in large part due to the uncertainties in the bottom-up emissions as discussed in Section
498 3.1.1. Satellite assimilation led to 2-6 ppbv higher RAQMS surface O₃ in the
499 central/southern/eastern US than in its free simulation, which are associated with higher positive
500 biases.

501

502 The surface O₃ performance by individual global models varies significantly, e.g., with the
503 RMSEs at all CASTNET sites ranging from ~9 ppbv to >15 ppbv (Table 2b). As reported in the
504 literature (e.g., Geddes et al., 2016; Travis et al., 2016), the representation of land use/land cover,
505 boundary layer mixing and chemistry can be sources of uncertainty for certain global model (i.e.,
506 GEOS-Chem), but how serious these issues were in the other models need to be investigated
507 further. Some other possible reasons include the variation of these models' non-anthropogenic
508 emission inputs and chemical mechanisms (Table 1c). Future work should emphasize on
509 evaluating and comparing all models on process level to better understand their performance.
510 Except in the northeastern US, the eight-model ensembles show better agreement with the
511 CASTNET O₃ observations than the three boundary condition-model ensemble. Overall the three-
512 model ensemble only outperforms one model but the eight-model ensemble outperforms seven
513 individuals. This reflects that averaging the results from a larger number of models in this case
514 more effectively cancelled out the positive or negative biases from the individual models.

515

516 The monthly-mean surface O₃ from multiple global models' free runs was also evaluated
517 with the EANET observations. Among the three boundary condition models, GEOS-Chem
518 produced higher O₃ than the other two throughout the year, and C-IFS O₃ is the lowest from April
519 to December. The three-model and eight-model ensembles are lower than the surface O₃
520 observations by <10 ppbv during high O₃ seasons (winter/spring), but show substantial (>10 ppbv)
521 positive biases during low O₃ seasons especially in July and August (Figure 3b), similar to the
522 HTAP1 model performance over Japan in 2001 (Fiore et al., 2009). During May-June 2010,
523 generally the models performed better at the Japanese sites than at the Korean sites (Table 2c),
524 with significant positive biases occurring at low O₃ regions (e.g., in central Japan) and negative

525 biases found at high O₃ regions, mainly owing to the uncertainty in the local and upwind emissions.
526 The different approaches to generate the monthly-mean modeled and the observed O₃ data may
527 have also contributed to these model-observation discrepancies. Overall O₃ performance by
528 individual models varies less significantly than at the CASTNET sites, with RMSEs ranging from
529 8.6 ppbv to ~13 ppbv (Table 2b). The three-model ensemble outperforms two individual models,
530 and the eight-model ensemble outperforms six individual models. Unlike at the CASTNET sites,
531 the three-model ensemble agrees better with the observations than the eight-model ensemble
532 (Table 2c).

533 3.1.3. Evaluation of the STEM regional base simulations w/ three sets of boundary conditions

536 The three STEM base simulations using different boundary conditions were evaluated with
537 the hourly O₃ observations at the CASTNET sites in the four US subregions. The evaluation
538 included the 8 May-30 June 2010 period to exclude the results during the one-week spin-up period.
539 The time series plots of observed and modeled O₃ at the western US CASTNET sites show that
540 STEM was capable of capturing several high O₃ periods, and it produced larger biases during the
541 nighttime (Figure 2c), as a result of the poorer WRF performance. Figure 2c and the evaluation
542 statistics in Table 3a-b indicate that STEM/C-IFS O₃ concentrations are associated with the highest
543 positive bias and RMSE, while the STEM/GEOS-Chem and STEM/RAQMS predictions were
544 positively and negatively biased by less than 2 ppbv, respectively, with similar RMSEs and
545 correlations with the observations. The quality of the three STEM simulation mean is closest to
546 the STEM/GEOS-Chem run, with the mean bias/RMSE of ~1.6/4.9 ppbv, much better than the
547 three-boundary model ensemble (-5.7/10.4 ppbv). However, this good performance can be a net
548 effect of incorrect partitioning between the trans-boundary and local source contributions, with the
549 former being underestimated and offsetting the overestimation of the latter. Switching the STEM
550 chemical boundary conditions to the assimilated RAQMS base simulation led to increases in the
551 simulated surface O₃ concentrations by >9 ppbv in the western US (Figure S2, right), associated
552 with higher positive biases (due to several factors discussed in the next paragraph). Regional-scale
553 assimilation could further reduce uncertainties introduced from regional meteorological and
554 emission inputs to obtain better modeled total O₃ and the partitioning of trans-boundary versus US
555 contributions (e.g., Huang et al., 2015).

557 The three STEM base simulations all significantly overpredicted O₃ over the rest of the US
558 in part due to the overall overestimated NO_x emissions, with the STEM/RAQMS associated with
559 the lowest RMSEs and mean biases, but STEM/C-IFS correlated best with the observations (Table
560 3b). These positive biases are higher than the global model ensembles', which can partially result
561 from the possible unrealistic VOC speciation of the emission inventory and the SAPRC 99
562 chemical mechanism: Although SAPRC mechanisms have been used in air quality modeling for
563 regulatory applications in some US states such as California, they usually produced higher O₃ than
564 other mechanisms such as the CB04 and the CB05 (which were used by some HTAP2 global
565 models, see Table 1c) over the US, and the comparisons between SAPRC 99 and SAPRC 2007
566 are still in progress (e.g., Luecken et al., 2008; Zhang et al., 2012; Cai et al., 2011). It is important
567 to timely update the chemical mechanisms in the chemistry models, and we also suggest to timely
568 upgrade the VOC speciation in the bottom-up emission inventories in the US to benefit the air
569 quality modeling. Additionally, the uncertainty from non-anthropogenic emissions, such as the
570 biogenic VOC emissions from WRF/MEGAN which is known to often have positive biases, can

571 be another cause: As Hogrefe et al. (2011) presented, the MEGAN emissions resulted in a higher
572 O₃ response to hypothetical anthropogenic NO_x emission reductions compared with another set of
573 biogenic emission input. Huang et al. (2017) showed that MEGAN's positive biases are in part
574 due to the positively-biased temperature and radiation in WRF, and reducing ~2°C in WRF's
575 temperature biases using a different land initialization approach led to ~20% decreases in
576 MEGAN's isoprene emission estimates in September 2013 over some southeastern US regions.
577 These temperature and radiation biases, can also be important sources of uncertainty in the
578 modeled O₃ production. Quantifying the impacts of overestimated biogenic emissions and the
579 biased weather fields that contributed to the biases in emissions on the modeled O₃ is still an
580 ongoing work. Some existing studies also reported O₃ and NO₂ biases from other regional models
581 in the eastern US, due to the chemical mechanism and biases in NO_x and biogenic VOC emissions
582 (e.g., Canty et al., 2015). We anticipate that the results from the Air Quality Model Evaluation
583 International Initiative (AQMEII) experiment (e.g., Schere et al., 2012; Solazzo et al., 2012;
584 Galmarini et al., 2015, 2017), which involves more regional model simulations over the US with
585 the similar set of boundary conditions but different chemical mechanisms and non-anthropogenic
586 emission inputs, can help better understand the causes of errors in the simulated total O₃.

587

588 3.2. *The NAM surface O₃ sensitivity to extra-regional anthropogenic pollutants*

589 3.2.1. Global model ensembles

590

591 The impact of all foreign (i.e. non-NAM) anthropogenic sources on NAM surface O₃ was
592 first explored, including the spatial distributions of the RERER metric (eq. (2)) based on various
593 global models' simulations (Figure 5), and the domain wide mean sensitivities R (O₃, non-NAM,
594 20%) (eq. (1d)) (Figure 6). Across the NAM, the strongest impacts were found in spring time
595 (March-April-May, larger than 1.5 ppbv in average over the domain) and the weakest impacts are
596 shown during the summertime (June-July-August, 1.0-1.3 ppbv), consistent with the existing
597 knowledge on the seasonal variability of the non-local pollution impacts on NAM for other years
598 (e.g., Fiore et al., 2009; Reidmiller et al., 2009). All global models indicate strong non-NAM
599 anthropogenic source impacts on the western US mainly due to the impact of its high elevation,
600 and also near the US-Mexico border areas, especially southern Texas, due to their vicinity to the
601 Mexican emission sources. Over the western states, stronger non-local impacts were reflected from
602 the results based on higher-horizontal resolution global models (e.g., the >0.6 RERER values from
603 the half degree EMEP model, corresponding to its higher R(O₃, non-NAM, 20%) values than the
604 other models'), similar to the findings in previous modeling studies (Lin et al., 2010, 2012a).
605 Although on a coarse horizontal resolution of 2.8°, OsloCTM3 suggests stronger extra-regional
606 source influences on the northwestern US and the US-Canada border regions than the other models.
607 Its largest number of vertical layers among all global models might be a cause. Larger-than-1
608 RERER values are often seen near the urban areas and large point sources due to the titration,
609 especially evident from the higher resolution model results. The R(O₃, EAS, 20%) values are larger
610 than 1/3 of the R(O₃, non-NAM, 20%) (0.2-0.5 ppbv from April to June), more than 3-4 times
611 higher than R(O₃, EUR, 20%) and R(O₃, SAS, 20%). Note that all eight models contributed to the
612 R(O₃, EAS, 20%) calculations, but one or two models did not provide all necessary sensitivity runs
613 to compute the RERER, R(O₃, non-NAM, 20%), R(O₃, EUR, 20%), or R(O₃, SAS, 20%).

614

615 Comparing to the HTAP1 modeling results, the magnitudes of R(O₃, EUR, 20%) from this
616 study are smaller by a factor of 2-3; In contrast, the R(O₃, non-NAM, 20%) and R(O₃, EAS, 20%)

617 values are >50% higher than the HTAP1 modeling results. The different HTAP1 and HTAP2
618 results are possibly due to the following three reasons: 1) the substantial improvement in the
619 European air quality over the past decades that is shown in Crippa et al. (2016) and Pouliot et al.
620 (2015), which contrasts with the growing anthropogenic emissions from the East Asia and other
621 developing countries during 2001-2010; 2) the changes in the HTAP2 experiment setup from
622 HTAP1. This includes the differences in the participating models, and the different region
623 definitions, e.g., EUR by HTAP1's definition includes regions in Russia/Belarus/Ukraine,
624 Middle East and North Africa that are excluded from the HTAP2's EUR domain. For EAS and
625 SAS, however, the regions not overlapped by HTAP1 and HTAP2 are mostly in the less
626 populated/polluted regions; 3) the stronger-than-normal transport in 2010 than in 2000-2001, as
627 first introduced in Section 2.2.1. Interannual variability of $R(O_3, \text{EAS}, 20\%)$ and $R(O_3, \text{non-NAM},$
628 $20\%)$ is also found between 2010 and 2008-2009, based on the SNU GEOS-Chem calculations
629 (Figure S3). Foreign anthropogenic pollution impact on NAM was stronger in 2010 than in 2008-
630 2009, especially in April-May. This can be in part due to the higher O_3 precursors' emissions in
631 2010 from extra-regions including the East Asia (Table S1), as well as the spring 2010
632 meteorological conditions that favored the trans-Pacific pollution transport.
633

634 These monthly- and regional-mean $R(O_3, \text{EAS}, 20\%)$ values suggest that despite dilution
635 along the great transport distance, the EAS anthropogenic sources still had distinguishable impact
636 on the NAM surface O_3 . Similar to the findings from the HTAP1 studies, the large intermodel
637 variability (as indicated in Table 4) in the estimates of intercontinental SR relationships indicates
638 the uncertainties of these models in representing the key atmospheric processes which needs more
639 investigations in the future. Figure 6b compares the $R(O_3, \text{EAS}, 20\%)$ estimated by individual
640 boundary condition models, their ensemble mean sensitivities, and the eight-global model mean.
641 The averaged $R(O_3, \text{EAS}, 20\%)$ from the boundary condition model results are smaller than the
642 eight-global model mean, and except for July-October 2010, GEOS-Chem gives higher $R(O_3, \text{EAS},$
643 $20\%)$ than RAQMS and C-IFS, consistent with its highest O_3 prediction in the EAS source region
644 (Figure 3b). Overall, $R(O_3, \text{EAS}, 20\%)$ and its intermodel differences are much smaller than the
645 biases of the modeled total O_3 in NAM. Other factors can contribute more significantly to the
646 biases in the modeled total O_3 , such as the stratospheric O_3 intrusion and the local O_3 formation,
647 and assessing the impacts from these factors would be also helpful for understanding the
648 uncertainties in the modeled O_3 .
649

650 The O_3 sensitivities in response to the perturbations of individual species or sector
651 emissions in East Asia, estimated by the GEOS-Chem adjoint model, were also analyzed (Figure
652 S3). These sensitivities show similar seasonal variability to $R(O_3, \text{EAS}, 20\%)$, with the values
653 ~twice as high in the spring than in summer, also consistent with the results on previous years
654 based on the 20% emission perturbation approach (e.g., Fiore et al., 2009; Brown-Steiner and Hess,
655 2011; Emmons et al., 2012). However, this seasonal variability is weaker than the results based on
656 the tagged tracer approach for earlier years: Using the CAM-Chem model, Brown-Steiner and
657 Hess (2011) reported that during the springtime, Asian O_3 created from the anthropogenic/biofuel
658 NO_x emissions affected NAM O_3 ~three times as strongly as in summer. This is because the
659 nonlinear O_3 chemistry, which is stronger outside of summer, caused larger O_3 responses to a 100%
660 reduction of NO_x emissions than 5 times of the O_3 responses to a 20% reduction of NO_x emissions.
661 The EAS anthropogenic NO_x emissions more strongly impacted the NAM surface O_3 than the
662 other major O_3 precursors, similar to the findings in Fiore et al. (2009) and Reidmiller et al. (2009)

663 using the perturbation approach, as well as the conclusions in Lapina et al. (2014) based on the
664 adjoint sensitivity analyses. Emissions from the power&industrial sectors are higher in East Asia
665 than the other sectors (Table S1), resulting in its stronger influences on the NAM surface O₃. As
666 the observed NO₂ columns started to drop since 2010 due to the effective denitration devices
667 implemented at the Chinese power and industrial plants (e.g., Liu et al., 2016), depending on the
668 changes in the VOC emissions, it is anticipated to see different R(O₃, EAS, 20%) values for the
669 years after 2010. Therefore, continued studies to assess the East Asian anthropogenic pollution
670 impacts on NAM during more recent years is needed. As emissions from various source sectors
671 can differ by their emitted altitudes and temporal (from diurnal to seasonal) profiles, efforts should
672 also be placed to have the models timely update the heights and temporal profiles of the emissions
673 from those various sectors.

674

675 3.2.2. Regional model sensitivities and their connections with the boundary condition models'

676

677 The monthly-mean STEM surface R(O₃, EAS, 20%) sensitivities based on different
678 boundary condition models were inter-compared, and also compared with the R(O₃, EAS, 20%)
679 estimated by their boundary condition models as well as the global model ensemble mean (Figure
680 7). For both May and June 2010, the domain-wide mean R(O₃, EAS, 20%) values from
681 STEM/RAQMS were higher than the estimates from RAQMS by 0.03 ppbv; the STEM/GEOS-
682 Chem R(O₃, EAS, 20%) values are lower than those of GEOS-Chem by 0.01-0.06 ppbv, and the
683 STEM/C-IFS R(O₃, EAS, 20%) is 0.02 ppbv higher than C-IFS's in June but slightly (<<0.01 ppbv)
684 lower in May. These differences are overall smaller than the inter-global model differences, and
685 can be due to various factors including the uncertainties in boundary condition chemical species
686 mapping, and the different meteorological/terrain fields/chemistry in the global and regional model
687 pairs. The STEM R(O₃, EAS, 20%) ensemble mean values, however, are less than 0.02 ppbv
688 different from its boundary condition model's ensemble mean for both months. The STEM R(O₃,
689 EAS, 20%) ensemble mean value in June is also close to the eight-global model ensemble mean,
690 but is ~0.05 ppbv lower than the eight-model mean in May. Choosing other/more global model
691 outputs as STEM's boundary conditions may lead to different STEM ensemble mean R(O₃, EAS,
692 20%) estimates. We also found that the period mean R(O₃, EAS, 20%) of ~0.2 ppbv sampled only
693 at the CASTNET sites (Table 3a) are smaller than those averaged in all model grids. This indicates
694 that currently the sparsely distributed surface network (especially over the western US that is more
695 strongly affected by the extra-regional sources than the other US regions) may miss many LRT
696 episodes that impact the NAM. The planned geostationary satellites with ~2-5 km footprint sizes
697 and hourly sampling frequency (Hilsenrath and Chance, 2013) will help better capture the high O₃
698 and LRT episodes in these regions.

699

700 The spatial patterns of the monthly-mean STEM surface R(O₃, EAS, 20%) sensitivities
701 based on the three boundary condition models are notably different, but overall resemble what's
702 estimated by the corresponding boundary condition model, and the STEM sensitivities show more
703 local details in certain high elevation regions in the US west (Figure 8 shows the June 2010
704 conditions as an example). These different sensitivities were investigated further, by examining
705 the R(O₃, EAS, 20%) values near the source regions (i.e., East Asia) as well as near the receptor
706 regions (Figure 9). More East Asian anthropogenic O₃ seems to be transported at the upper
707 troposphere in RAQMS than in the other two models. GEOS-Chem and RAQMS R(O₃, EAS, 20%)
708 sensitivities are similar over the EAS as well as the 500-900 hPa near the receptor in the eastern

709 Pacific (at $\sim 135^\circ\text{W}$), the altitudes US surface O_3 are most strongly sensitive to during the
710 summertime as concluded from previous studies (e.g., Huang et al., 2010, 2013a; Parrish et al.,
711 2010). Despite the close NAM domain-wide mean values from the STEM/GEOS-Chem and
712 STEM/RAQMS, the spatial patterns of $R(\text{O}_3, \text{EAS}, 20\%)$ over NAM differ in these two cases,
713 with the latter case showing sharper gradients especially in the western US, partially due to the
714 impact of its higher horizontal resolution. The $R(\text{O}_3, \text{EAS}, 20\%)$ values from STEM/C-IFS are
715 lower than from the other two cases both near the sources and at (near) NAM. The STEM surface
716 (also near surface, not shown in figures) $R(\text{O}_3, \text{EAS}, 20\%)$ does not spatially correlate well with
717 the column $R(\text{O}_3, \text{EAS}, 20\%)$, the latter of which contributed more to the base case O_3 columns,
718 indicating that a good portion of the transported East Asian pollution did not descend to the lower
719 altitudes to impact the boundary layer/ground level air quality. An additional regional simulation
720 was performed in which the STEM boundary conditions were downscaled from a RAQMS
721 simulation without the East Asian anthropogenic emissions. The non-linear emission perturbation-
722 O_3 response relationships, as the larger-than-1 S_{O_3} metric (eq. (3)) indicate, are seen across the
723 domain, for both the surface and column O_3 (Figure 8). S_{O_3} for column O_3 , ranging from 1.15-1.25
724 in most regions, are overall ~ 0.05 higher than S_{O_3} for the surface O_3 . Therefore, the full source
725 contribution obtained by linearly scaling the receptor regional mean O_3 sensitivity to the 20%
726 reduction in the source region emissions may be underestimated by at least $\sim 10\%$.

727

728 3.2.3. Regional model MDA8 sensitivities on all days and during the O_3 exceedances

729 The temporal variability of the STEM $R(\text{O}_3, \text{EAS}, 20\%)$ ensemble sensitivities were also
730 studied. For most US subregions, 3-6 LRT episodes (defined as when the sensitivities are above
731 the period mean) were identified during May-June. Throughout this period, the hourly $R(\text{O}_3, \text{EAS},$
732 $20\%)$ and the observed O_3 at the surface CASTNET sites are weakly correlated (Table 3a), but
733 they display similar diurnal cycles (e.g., Figures 2c and 2d for the western US sites), possibly
734 because the deeper boundary layer depth during the daytime enhanced entrainment down-mixing
735 of the extra-regional pollutants to the surface. The identified diurnal variability of the $R(\text{O}_3, \text{EAS},$
736 $20\%)$ can cause differences in the calculated MDA8 and all-hour mean $R(\text{O}_3, \text{EAS}, 20\%)$ values.
737 Figure S4 shows that the mean $R(\text{MDA8}, \text{EAS}, 20\%)$ values, usually at daytimes, are higher than
738 the all-hour averaged $R(\text{O}_3, \text{EAS}, 20\%)$ in most STEM model grids during both months. Therefore,
739 it is important for more HTAP2 participating models to save their outputs hourly in order to
740 conveniently compute the policy-relevant metrics for the O_3 sensitivities. Also, the hourly
741 sampling frequency of the planned geostationary satellites is anticipated to be more helpful for
742 evaluating the impacts of the LRT episodes.

743

744 The STEM $R(\text{MDA8}, \text{EAS}, 20\%)$ in all model grids within the four US subregions were
745 averaged on all days during May-June 2010 and only on the days when the simulated total MDA8
746 O_3 is over 70 ppbv (Figure 10). These sensitivities also show appreciable spatial variability: from
747 0.35-0.58 ppbv in the western US (also with the largest standard deviations, not shown), which is
748 slightly higher than the HTAP1 results reported by Reidmiller et al. (2009) for Spring 2001, to
749 ~ 0.1 -0.25 ppbv in the rest three subregions, which is close to the Reidmiller et al. (2009) results.

750

751 Comparing the solid bar plots in Figures 10-11, we found that on all days in the three non-
752 western subregions, $R(\text{MDA8}, \text{EAS}, 20\%)$ values sampled at CASTNET sites are slightly smaller
753 than those computed for all model grids, while in the non-western states the opposite differences

754 are seen. This again suggests that expanding observation network would help better capture the
755 high O₃ and LRT episodes.

756
757 Figure 10 suggests smaller R(MDA8, EAS, 20%) values during the high O₃ days in all
758 subregions. However, STEM's total O₃ concentrations at CASTNET sites during the O₃
759 exceedances were substantially overpredicted in non-western US regions while significantly
760 underpredicted in the western US (see mean biases above the bar plots in Figure 11). Therefore,
761 the R(MDA8, EAS, 20%) values shown in Figure 10 during O₃ exceedances can actually represent
762 the sensitivities during the non-exceedances in non-western US regions, and may not represent the
763 sensitivities during all O₃ exceedances in the western US. Figures 11-12 show that if calculated
764 only at the CASTNET sites during the exceedances, in non-western US regions, R(MDA8, EAS,
765 20%) is 0.02-0.07 ppbv smaller during the high O₃ total days. This is qualitatively consistent with
766 the findings in Reidmiller et al. (2009), and is possibly because that the LRT impacts were stronger
767 on some days with good dispersion conditions when the NAAQS was not exceeded, but weaker
768 on some high O₃ days under stagnant conditions. In contrast, western US R(MDA8, EAS, 20%) at
769 CASTNET sites was ~0.05 ppbv higher on high O₃ days than for all days, and this differences are
770 larger in rural/remote areas where local influences are less dominant. As a result, the
771 medium/strong positive correlations are found between modeled LRT of pollution and the total O₃
772 in these regions (Table 3a; Lin et al., 2012a).

773
774 3.3. *Case studies of spring (9 May) and summer (10 June) LRT events mixed with stratospheric*
775 *O₃ intrusions*

776
777 Lin et al. (2012a, b) and Neuman et al. (2012) showed that the trans-Pacific pollution
778 transport intensely impacted the western US during 8-10 May, 2010, intermingled with a
779 stratospheric intrusion that contributed to at least 1/3 of the total O₃ in some high elevation regions.
780 This episode is indeed indicated by the O₃ and CO products from AIRS and TES at ~500 hPa over
781 the Eastern Pacific (Figure 13), and the observed TES and IASI O₃ profiles over the western US
782 indicated elevated O₃ levels (>80 ppbv) at 700-900 hPa. Huang et al. (2013b) found that the
783 meteorological conditions during this period (i.e., a strong jet at ~700 hPa with wind speed >20
784 m/s shifted southwesterly when passing the southern California and continued to travel towards
785 the mountain states), along with the orographic lifting, efficiently exported the southern California
786 anthropogenic pollution, which was chemically coupled with the extra-regional pollution and
787 significantly enhanced the O₃ levels in the US intermountain west.

788
789 We selected this episode to compare the STEM surface total O₃ concentrations as well as
790 the R(O₃, EAS, 20%) sensitivities based on the different HTAP2 boundary condition models.
791 Figure 14 evaluates the simulated O₃ profiles in the western US from several STEM base
792 simulations against the TES and IASI O₃ retrievals, and Figures 15a-d indicate the performance of
793 the daily surface total MDA8 O₃ from these simulations. We found that the underestimated free
794 tropospheric O₃ from the STEM simulations that used any single free-running chemical boundary
795 conditions contributed to the underestimated STEM surface O₃ in the high elevation mountain
796 states: e.g., by 9-14 ppbv at three CASTNET sites (Grand Canyon National Park (NP), AZ;
797 Canyonlands NP, UT; and Rocky Mountain NP, CO) where O₃ exceedances were observed. The
798 unsatisfactory performance by free-running global models during high O₃ events would pose
799 difficulties for regional models (regardless of their resolutions and other configurations,

800 parameterization) to accurately estimate the SR relationships using boundary conditions
801 downscaled from these model runs. The STEM base simulation using the RAQMS assimilated
802 fields as the boundary conditions, agrees most with the observed O₃ at the CASTNET sites, as well
803 as the TES and IASI O₃ profiles in the western states. Similar to the conclusions drawn in Huang
804 et al. (2010, 2015) for summer 2008, we again demonstrated the robustness of satellite chemical
805 data assimilation for improving the boundary condition models' O₃ performance. As the
806 enhancement of O₃ due to the assimilation is much larger than the O₃ sensitivities to the EAS
807 anthropogenic emissions, the assimilation mainly improved the contributions from other sources,
808 such as the stratospheric O₃.

809
810 The quality of the model boundary conditions only indicates how well the total “transported
811 background” component is represented, and can not be directly connected with the accuracy of the
812 model estimated R(O₃, EAS, 20%) sensitivities, which also show notable intermodel differences:
813 The estimated R(MDA8, EAS, 20%) in the different STEM cases range from <1.0 ppbv to ~1.3
814 ppbv, at least 40% higher than the May-June period mean in Figures 10-11. The mean R(MDA8,
815 EAS, 20%) at three high O₃ CASTNET sites range from 0.73 (STEM/GEOS-Chem) to 0.98 ppbv
816 (STEM/C-IFS), with the mean S_{O₃} of ~1.14 at these sites based on the STEM/RAQMS runs due
817 to the nonlinear emission perturbation-O₃ response relationships (Figure 15e-h). The R(MDA8,
818 EAS, 100%) from the STEM/RAQMS case is as high as >7 ppbv over the high terrain regions.
819 These are of smaller magnitudes than the estimates in Lin et al. (2012a), possibly due to the
820 differences in the used models and the bottom-up emission inputs.

821
822 A stratospheric O₃ intrusion also affected the NE US on the same day, as revealed by the
823 satellite free tropospheric O₃ and CO observations (Figure 13). This intrusion was mixed with LRT
824 East Asian pollution (Figure 15 and Figure S5). However, this intrusion did not enhance the NE
825 boundary layer/surface O₃ concentrations, which were actually anomalously low (MDA8<40 ppbv)
826 as indicated by the model base simulations and the CASTNET observations (Figure 15a-d).
827 Similar characteristics during summertime stratospheric O₃ intrusion events around this region
828 have been discussed by Ott et al. (2016). The East Asian pollution less intensely (<50%) affected
829 the surface O₃ levels in these regions than in the US west, due to the greater transport distances,
830 stronger local emission influence on chemical production/loss, as well as the impact of the overall
831 flat terrain in the US east.

832
833 A summertime LRT event on 9-10 June is analyzed to contrast with the 9 May case study.
834 Lin et al. (2012b) showed that >80 ppbv of ozonesonde data in northern California at 2-6 km
835 measured the stratospheric O₃ remnants during this episode, and the transported stratospheric O₃
836 contributed to as much as ~50% of the total O₃ in southern California based on their model
837 calculations. We show that on 10 June over 100 ppbv of O₃, as well as <90 ppbv CO, was observed
838 by satellites at ~500 hPa above Nevada and northern California (Figure 16), which again was
839 substantially underestimated by all free-running models (Figure 17), resulting in the
840 underpredicted total O₃ at two CASTNET sites in southern California (Converse Station and
841 Joshua Tree NP) that experienced O₃ exceedances on this day (Figure 18a-c). The negative biases
842 in the “transported background” O₃ and surface MDA8 O₃ were successfully reduced by
843 incorporating satellite data (Figures 17 and 18d).

844

845 Figures 18e-h show that LRT of EAS anthropogenic pollution also strongly affected
846 southern California and Nevada. Notable intermodel differences are again found in the estimated
847 R(MDA8, EAS, 20%), but they are overall lower than on 9 May (<1.0 ppbv). The mean R(MDA8,
848 EAS, 20%) at the two high O₃ CASTNET sites range from 0.54 (STEM/C-IFS) to 0.86 ppbv
849 (STEM/RAQMS), with the mean S_{O₃} of ~1.13 at these sites based on the STEM/RAQMS runs
850 (Figure 18e-h). The R(MDA8, EAS, 100%) from the STEM/RAQMS case is as high as >6 ppbv
851 over southern California and Nevada. Compared to the spring event, R(MDA8, EAS, 20%) in the
852 eastern US are discernable only over a limited region, due to weaker transport and stronger local
853 chemical production/loss.

854

855 **4. Conclusions and suggestions on future directions**

856

857 In support of the HTAP Phase 2 experiment that involved high-resolution global models
858 and regional models' participation to advance the understanding of the pollutants' SR relationships
859 in the Northern Hemisphere, we conducted a number of regional scale STEM base and forward
860 sensitivity simulations over NAM during May-June 2010. STEM's top and lateral chemical
861 boundary conditions were downscaled from three global models' (i.e., GEOS-Chem, RAQMS,
862 and ECMWF C-IFS) base and sensitivity simulations (in which the East Asian anthropogenic
863 emissions were reduced by 20%). Despite dilution along the great transport distance, the East
864 Asian anthropogenic sources still had distinguishable impact on the NAM surface O₃, with the
865 period-mean NAM O₃ sensitivities to a 20% reduction of the East Asian anthropogenic emissions
866 (i.e., R(O₃, EAS, 20%)) ranging from ~0.24 ppbv (STEM/C-IFS) to ~0.34 ppbv (STEM/RAQMS).
867 The spatial patterns of the STEM surface O₃ sensitivities over NAM overall resembled those from
868 its corresponding boundary condition model, with regional/period mean R(O₃, EAS, 20%) differed
869 slightly (<10%) from its corresponding boundary condition model's, which are smaller than those
870 among its boundary condition models. The boundary condition models' two-month mean R(O₃,
871 EAS, 20%) was ~8% lower than the mean sensitivity estimated by multiple global models.
872 Therefore, choosing other global model outputs as STEM's boundary conditions may lead to
873 different STEM O₃ sensitivities. The biases and RMSEs in the simulated total O₃, which differed
874 significantly among models, can partially be due to the uncertainty in the bottom-up NO_x emission
875 inputs according to the model comparison with the OMI NO₂ columns, and future work on
876 attributing the intermodel differences on process level is particularly important for better
877 understanding the sources of uncertainties in the modeled total O₃ and its source contribution.

878

879 The HTAP2 multi-model ensemble mean R(O₃, EAS, 20%) values in 2010 were higher
880 than the HTAP1 reported 2001 conditions, due to the impacts of the growing East Asian
881 anthropogenic emissions, the interannual variability in atmospheric circulation (i.e., stronger trans-
882 Pacific transport in spring 2010 following an El Niño event), and the different experiment designs
883 of HTAP1 and HTAP2. The GEOS-Chem O₃ sensitivities in 2010 were also higher than the 2008-
884 2009 conditions due to the increasing Asian emissions and the spring 2010 meteorological
885 conditions that favored the trans-Pacific pollution transport. The GEOS-Chem sensitivity
886 calculations indicate that the East Asian anthropogenic NO_x emissions mattered more than the
887 other East Asian O₃ precursors to the NAM O₃, qualitatively consistent with previous adjoint
888 sensitivity calculations. Continued research is needed on temporal changes of emissions for
889 different species and sectors in NAM and foreign countries as well as their impacts on the SR
890 relationships. As emissions from various source sectors can differ by emitted altitudes and

891 temporal profiles, efforts should also be placed to have the models timely update the height and
892 temporal profiles of the emissions from various sectors.

893
894 An additional STEM simulation was performed in which the boundary conditions were
895 downscaled from a RAQMS simulation without East Asian anthropogenic emissions (i.e., a 100%
896 emission reduction), to assess the scalability of the mean O₃ sensitivities to the size of the emission
897 perturbation. The scalability was found to be spatially varying, ranging from 1.15-1.25 for column
898 O₃ in most US regions, which were overall ~0.05 higher than the surface O₃'s. Therefore, the full
899 source contribution obtained by linearly scaling the NAM regional mean O₃ sensitivity to the 20%
900 reduction in the East Asian emissions may be underestimated by at least 10%. The underestimation
901 in other seasons of the HTAP2 study period may be higher and will need to be quantified in future
902 work. Also, motivated by Lapina et al. (2014), additional calculations will be conducted in future
903 to explore the scalability of different O₃ metrics in these cases. For future source attribution
904 analysis, in general it is recommended to directly choose the suitable size of the emission
905 perturbation based on the specific questions to address, and to avoid linearly scaling O₃
906 sensitivities that are based on other amounts of the perturbations.

907
908 The STEM O₃ sensitivities to the East Asian anthropogenic emissions (based on three
909 boundary condition models separately and averagely) were strong during 3-6 episodes in May-
910 June 2010, following similar diurnal cycles as the total O₃. Stronger-than-normal East Asian
911 anthropogenic pollution impacts were estimated during O₃ exceedances in the western US,
912 especially over the high terrain rural/remote areas; in contrast, non-local pollution impacts were
913 less strong during O₃ exceedances in other US regions. We emphasized the importance of saving
914 model results hourly for continentally calculate policy-relevant metrics, as well as the usefulness of
915 hourly sampling frequency of the planned geostationary satellites for better evaluating the impacts
916 of the LRT events.

917
918 Based on model calculations, satellite O₃ (TES, JPL-IASI, and AIRS), CO (TES and AIRS)
919 and surface O₃ observations on 9 May 2010, we showed the different influences from stratospheric
920 O₃ intrusions along with the transported East Asian pollution on O₃ in the western and the eastern
921 US. This event was further compared with a summer event of 10 June 2010. During both events,
922 the unsatisfactory performance of free-running global models would pose difficulties for regional
923 models (regardless of their resolutions and other configurations, parameterization) to accurately
924 simulate the surface O₃ and its source contribution using boundary conditions downscaled from
925 these model runs. Incorporating satellite (OMI and MLS) O₃ data effectively improved the
926 modeled O₃. As chemical data assimilation techniques keep developing (Bocquet et al., 2015),
927 several HTAP2 participating global models have already been able to assimilate single- or multi-
928 constitute satellite atmospheric composition data (e.g., Miyazaki et al., 2012; Parrington et al.,
929 2008, 2009; Huang et al., 2015; Inness et al., 2015; Flemming et al., 2017). Comparing the
930 performance of the assimilated fields from different models, and making the global model
931 assimilated chemical fields in the suitable format for being used as boundary conditions would be
932 very beneficial for future regional modeling, as well as for better interpreting the pollutants'
933 distributions especially during the exceptional events. Meanwhile, efforts should also be devoted
934 to advancing and applying higher-resolution regional scale modeling and chemical data
935 assimilation. Furthermore, although satellite observations have been applied for improving the
936 estimated US background O₃ (e.g., Huang et al., 2015), using satellite (and/or other types of)

937 observations to improve SR relationship studies also needs to be explored. Some of the possible
938 methods include: 1) The combination of data assimilation and the tagging approach; 2) Introducing
939 observation-constrained emission estimates in the emission perturbation analyses.

940

941 **Acknowledgements**

942

943 The global and regional modeling results used in this study have been submitted to the
944 AeroCom database following the HTAP2 data submission guidelines ([http://iek8wikis.iek.fz-](http://iek8wikis.iek.fz-juelich.de/HTAPWiki/HTAP-2-data-submission)
945 [juelich.de/HTAPWiki/HTAP-2-data-submission](http://iek8wikis.iek.fz-juelich.de/HTAPWiki/HTAP-2-data-submission)), or can be made available upon request.
946 Technical support from Anna Carlin Benedictow, Brigitte Koffi, Jan Griesfeller, and Michael
947 Schulz regarding formatting and submitting the data to the AeroCom is acknowledged. MH thanks
948 the research resources at the University of Iowa and JPL/Caltech that supported this study, as well
949 as the travel funding from the US EPA for attending the related HTAP2 workshops. DKH and YD
950 recognize support from NASA AQUEST. FD Acknowledges support from the Administrative
951 Arrangement. Part of this research was carried out at the Jet Propulsion Laboratory, California
952 Institute of Technology, under contract to the National Aeronautics and Space Administration.
953 Reference herein to any specific commercial product, process or service by trade name, trademark,
954 manufacturer or otherwise does not constitute or imply its endorsement by the United States
955 Government or the Jet Propulsion Laboratory, California Institute of Technology. The views,
956 opinions, and findings contained in this report are those of the author(s) and should not be
957 construed as an official National Oceanic and Atmospheric Administration or U.S. Government
958 position, policy, or decision. We also acknowledge the feedbacks from Dr. Gail Tonnesen, two
959 anonymous reviewers, and Dr. Meiyun Lin on earlier versions of this paper, that helped improve
960 its quality.

961

962 **References**

963

- 964 Anderson, D. C., Loughner, C. P., Diskin, G., Weinheimer, A., Canty, T., P., Salawitch, R. J.,
965 Worden, H. M., Fried, A., 25 Mikoviny, T., Wisthaler, A., and Dickerson, R., R. (2014),
966 Measured and modeled CO and NO_y in DISCOVER-AQ: An evaluation of emissions and
967 chemistry over the eastern US, *Atmos. Environ.*, 96, 78-87, doi:
968 10.1016/j.atmosenv.2014.07.004.
- 969 Allen, D. J., Pickering, K. E., Pinder, R. W., Henderson, B. H., Appel, K. W., and Prados, A.
970 (2012), Impact of lightning-NO on eastern United States photochemistry during the summer
971 of 2006 as determined using the CMAQ model, *Atmos. Chem. Phys.*, 12, 1737-1758, doi:
972 10.5194/acp-12-1737-2012.
- 973 Ambrose, J.L., Reidmiller, D.R., and Jaffe, D.A. (2011), Causes of high O₃ in the lower free
974 troposphere over the Pacific Northwest as observed at the Mt. Bachelor Observatory. *Atmos.*
975 *Environ.*, 45, 5302–5315, doi: 10.1016/j.atmosenv.2011.06.056.
- 976 Anenberg, S. C., L. W. Horowitz, D. Q. Tong, and J. J. West (2010), An estimate of the global
977 burden of anthropogenic ozone and fine particulate matter on premature human mortality using
978 atmospheric modeling, *Environ. Health Perspect.*, 118(9), 1189–1195.
- 979 Avnery, S, D.L. Mauzerall, J. Liu, L.W. Horowitz (2011a), Global Crop Yield Reductions due to
980 Surface Ozone Exposure: 1. Year 2000 Crop Production Losses and Economic
981 Damage, *Atmos. Environ.*, 45, 2284-2296.

982 Avnery, S, D.L. Mauzerall, J. Liu, L.W. Horowitz (2011b), Global Crop Yield Reductions due to
983 Surface Ozone Exposure: 2. Year 2030 Potential Crop Production Losses and Economic
984 Damage under Two Scenarios of O₃ Pollution, *Atmos. Environ.*, 45, 2297-2309.

985 Beer, R., T. A. Glavich, and D. M. Rider (2001), Tropospheric emission spectrometer for the Earth
986 Observing System's Aura satellite, *Applied Optics*, 40, 2356 – 2367.

987 Beer, R (2006), TES on the Aura Mission: Scientific Objectives, Measurements, and Analysis
988 Overview, *IEEE Transaction on Geoscience and Remote Sensing*, 44, 1102-1105.

989 Bian, J., A. Gettelman, H. Chen, and L. L. Pan (2007), Validation of satellite ozone profile
990 retrievals using Beijing ozonesonde data, *J. Geophys. Res.*, 112, D06305,
991 doi:10.1029/2006JD007502.

992 Bocquet, M., Elbern, H., Eskes, H., Hirtl, M., Žabkar, R., Carmichael, G. R., Flemming, J., Inness,
993 A., Pagowski, M., Pérez Camaño, J. L., Saide, P. E., San Jose, R., Sofiev, M., Vira, J.,
994 Baklanov, A., Carnevale, C., Grell, G., and Seigneur, C. (2015), Data assimilation in
995 atmospheric chemistry models: current status and future prospects for coupled chemistry
996 meteorology models, *Atmos. Chem. Phys.*, 15, 5325-5358, doi:10.5194/acp-15-5325-2015.

997 Boersma, K. F., Braak, R., van der A, R. J. (2011a), Dutch OMI NO₂ (DOMINO) data product
998 v2.0 HE5 data file user manual. http://www.temis.nl/docs/OMI_NO2_HE5_2.0_2011.pdf.

999 Boersma, K. F., Eskes, H. J., Dirksen, R. J., van der A, R. J., Veefkind, J. P., Stammes, P., Huijnen,
1000 V., Kleipool, Q. L., Sneep, M., Claas, J., Leitão, J., Richter, A., Zhou, Y., Brunner, D. (2011b),
1001 An improved tropospheric NO₂ column retrieval algorithm for the Ozone Monitoring
1002 Instrument, *Atmos. Meas. Tech.*, 4, 1905-1928.

1003 Bowman, K. W., Rodgers, C. D., Kulawik, S. S., Worden, J., Sarkissian, E., Osterman, G., Steck,
1004 T., Lou, M., Eldering, A., Shephard, M., Worden, H., Lampel, M., Clough, S., Brown, P.,
1005 Rinsland, C., Gunson, M., and Beer, R. (2006), Tropospheric Emission Spectrometer:
1006 Retrieval method and error analysis, *IEEE Transaction on Geoscience and Remote Sensing*,
1007 44 (5), 1297–1307, doi: 10.1109/TGRS.2006.871234.

1008 Bowman, K., and D. K. Henze (2012), Attribution of direct ozone radiative forcing to spatially
1009 resolved emissions, *Geophys. Res. Lett.*, 39, L22704, doi:10.1029/2012GL053274.

1010 Brioude, J., Angevine, W. M., Ahmadov, R., Kim, S.-W., Evan, S., McKeen, S. A., Hsie, E.-Y.,
1011 Frost, G. J., Neuman, J. A., Pollack, I. B., Peischl, J., Ryerson, T. B., Holloway, J., Brown, S.
1012 S., Nowak, J. B., Roberts, J. M., Wofsy, S. C., Santoni, G. W., Oda, T., and Trainer, M. (2013),
1013 Top-down estimate of surface flux in the Los Angeles Basin using a mesoscale inverse
1014 modeling technique: assessing anthropogenic emissions of CO, NO_x and CO₂ and their
1015 impacts, *Atmos. Chem. Phys.*, 13, 3661-3677, doi:10.5194/acp-13-3661-2013.

1016 Brown-Steiner, B., and P. Hess (2011), Asian influence on surface ozone in the United States: A
1017 comparison of chemistry, seasonality, and transport mechanisms, *J. Geophys. Res.*, 116,
1018 D17309, doi:10.1029/2011JD015846.

1019 Cai, C., J. T. Kelly, J. C. Avise, A. P. Kaduwela, and W. R. Stockwell (2011), Photochemical
1020 Modeling in California with Two Chemical Mechanisms: Model Intercomparison and
1021 Response to Emission Reductions, *J. Air & Waste Manage. Assoc.*, 61:5, 559-572, doi:
1022 10.3155/1047-3289.61.5.559.

1023 Canty, T. P., Hembeck, L., Vinciguerra, T. P., Anderson, D. C., Goldberg, D. L., Carpenter, S. F.,
1024 Allen, D. J., Loughner, C. P., Salawitch, R. J., and Dickerson, R. R. (2015), Ozone and NO_x
1025 chemistry in the eastern US: evaluation of CMAQ/CB05 with satellite (OMI) data, *Atmos.*
1026 *Chem. Phys.*, 15, 10965-10982, doi:10.5194/acp-15-10965-2015.

1027 Carmichael, G.R., Tang, Y., Kurata, G., Uno, I., Streets, D.G., Thongboonchoo, N., Woo, J.H.,
1028 Guttikunda, S., White, A., Wang, T., Blake, D.R., Atlas, E., Fried, A., Potter, B., Avery, M.A.,
1029 Sachse, G.W., Sandholm, S.T., Kondo, Y., Talbot, R.W., Bandy, A., Thornton, D., and Clarke,
1030 A.D. (2003a), Evaluating regional emission estimates using the TRACE-P observations, *J.*
1031 *Geophys. Res.*, 108 (D21), 8810, doi: 10.1029/2002JD003116.

1032 Carmichael, G.R., Tang, Y., Kurata, G., Uno, I., Streets, D., Woo, J.H., Huang, H., Yienger, J.,
1033 Lefer, B., Shetter, R., Blake, D., Atlas, E., Fried, A., Apel, E., Eisele, F., Cantrell, C., Avery,
1034 M., Barrick, J., Sachse, G., Brune, W., Sandholm, S., Kondo, Y., Singh, H., Talbot, R., Bandy,
1035 A., Thornton, D., Clarke, A., and Heikes, B. (2003b), Regional-scale chemical transport
1036 modeling in support of the analysis of observations obtained during the TRACE-P experiment,
1037 *J. Geophys. Res.*, 108 (D21), 8823, doi: 10.1029/2002JD003117.

1038 Carter, W. P. L. (2000), Documentation of the SAPRC-99 chemical mechanism for VOC
1039 Reactivity Assessment, final report to California Air Resources Board, Contract No. 92-329
1040 and 95-308.

1041 Cooper, O. R., et al. (2010), Increasing springtime ozone mixing ratios in the free troposphere over
1042 western North America, *Nature*, 463, doi: 10.1038/nature08708.

1043 Cooper, O. R., Oltmans, S. J., Johnson, B. J., Brioude, J., Angevine, W., Trainer, M., Parrish, D.
1044 D., Ryerson, T. R., Pollack, I., Cullis, P. D., Ives, M. A., Tarasick, D. W., Al-Saadi, J., and
1045 Stajner, I. (2011), Measurement of western U.S. baseline ozone from the surface to the
1046 tropopause and assessment of downwind impact regions, *J. Geophys. Res.*, 116, D00V03, doi:
1047 10.1029/2011JD016095.

1048 Cooper, O., et al. (2016), Western NA Performance Evaluation for HTAP2, HTAP2 workshop,
1049 Potsdam, Germany, 2016.

1050 Crippa, M., Janssens-Maenhout, G., Dentener, F., Guizzardi, D., Sindelarova, K., Muntean, M.,
1051 Van Dingenen, R., and Granier, C. (2016), Forty years of improvements in European air
1052 quality: regional policy-industry interactions with global impacts, *Atmos. Chem. Phys.*, 16,
1053 3825-3841, doi:10.5194/acp-16-3825-2016.

1054 Emmons, L. K., Hess, P. G., Lamarque, J.-F., and Pfister, G. G. (2012), Tagged ozone mechanism
1055 for MOZART-4, CAM-chem and other chemical transport models, *Geosci. Model Dev.*, 5,
1056 1531-1542, doi:10.5194/gmd-5-1531-2012.

1057 Eskes, H. J. and Boersma, K. F. (2003), Averaging kernels for DOAS total-column satellite
1058 retrievals, *Atmos. Chem. Phys.*, 3, 1285-1291.

1059 Fiore, A. M., et al. (2009), Multimodel estimates of intercontinental source receptor relationships
1060 for ozone pollution, *J. Geophys. Res.*, 114, D04301, doi:10.1029/2008JD010816.

1061 Fiore, A. M., J. T. Oberman, M. Y. Lin, L. Zhang, O. E. Clifton, D. J. Jacob, V. Naik, L. W.
1062 Horowitz, J. P. Pinto, and G. P. Milly (2014), Estimating North American background ozone
1063 in U.S. surface air with two independent global models: Variability, uncertainties, and
1064 recommendations, *Atmos. Environ.*, 96, 284-300, doi: 10.1016/j.atmosenv.2014.07.045.

1065 Flemming, J., Huijnen, V., Arteta, J., Bechtold, P., Beljaars, A., Blechschmidt, A.-M., Diamantakis,
1066 M., Engelen, R. J., Gaudel, A., Inness, A., Jones, L., Josse, B., Katragkou, E., Marecal, V.,
1067 Peuch, V.-H., Richter, A., Schultz, M. G., Stein, O., and Tsikerdekis, A. (2015), Tropospheric
1068 chemistry in the Integrated Forecasting System of ECMWF, *Geosci. Model Dev.*, 8, 975-1003,
1069 doi:10.5194/gmd-8-975-2015.

1070 Flemming, J., Benedetti, A., Inness, A., Engelen, R., Jones, L., Huijnen, V., Remy, S., Parrington,
1071 M., Suttie, M., Bozzo, A., Peuch, V.-H., Akritidis, D., and Katragkou, E. (2017), The CAMS

1072 interim Reanalysis of Carbon Monoxide, Ozone and Aerosol for 2003–2015, *Atmos. Chem.*
1073 *Phys.*, 17, 1945–1983, doi:10.5194/acp-17-1945-2017.

1074 Galmarini, S., C. Hogrefe, D. Brunner, P. Makar, A. Baklanov (2015), Preface to the AQMEII p2
1075 Special issue, *Atmos. Environ.*, 115, 340–344.

1076 Galmarini, S., Koffi, B., Solazzo, E., Keating, T., Hogrefe, C., Schulz, M., Benedictow, A.,
1077 Griesfeller, J. J., Janssens-Maenhout, G., Carmichael, G., Fu, J., and Dentener, F. (2017),
1078 Technical note: Coordination and harmonization of the multi-scale, multi-model activities
1079 HTAP2, AQMEII3, and MICS-Asia3: simulations, emission inventories, boundary conditions,
1080 and model output formats, *Atmos. Chem. Phys.*, 17, 1543–1555, doi:10.5194/acp-17-1543-
1081 2017.

1082 Geddes, J. A., Heald, C. L., Silva, S. J., and Martin, R. V. (2016), Land cover change impacts on
1083 atmospheric chemistry: simulating projected large-scale tree mortality in the United States,
1084 *Atmos. Chem. Phys.*, 16, 2323–2340, doi:10.5194/acp-16-2323-2016.

1085 Gery, M. W., G. Z. Whitten, J. P. Killus, and M. C. Dodge (1989), A photochemical kinetics
1086 mechanism for urban and regional scale computer modeling, *J. Geophys. Res.*, 94, 12,925 –
1087 12,956, doi:10.1029/JD094iD10p12925.

1088 Granier, C., Lamarque, J. F., Mieville, A., Muller, J. F., Olivier, J., Orlando, J., Peters, J., Petron,
1089 G., Tyndall, G., and Wallens, S. (2005), POET, a database of surface emissions of ozone
1090 precursors, <http://www.aero.jussieu.fr/projet/ACCENT/POET.php>.

1091 Gratz, L.E., Jaffe, D.A., and Hee, J.R. (2014), Causes of increasing ozone and decreasing carbon
1092 monoxide in springtime at the Mt. Bachelor Observatory from 2004 to 2013, *Atmos. Environ.*,
1093 109, 323–330, doi: 10.1016/j.atmosenv.2014.05.076.

1094 Guenther, A. B., X. Jiang, C. L. Heald, T. Sakulyanontvittaya, T. Duhl, L. K. Emmons, and X.
1095 Wang (2012), The Model of Emissions of Gases and Aerosols from Nature version 2.1
1096 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions, *Geosci.*
1097 *Model Dev.*, 5 (6), 1471–1492.

1098 Henze, D. K., Hakami, A., and Seinfeld, J. H. (2007), Development of the adjoint of GEOS-Chem,
1099 *Atmos. Chem. Phys.*, 7, 2413–2433, doi:10.5194/acp-7-2413-2007.

1100 Hilsenrath, E., and K. Chance (2013), NASA ups the TEMPO on monitoring air pollution, *Earth*
1101 *Obs.*, 25, 10–15.

1102 Hogrefe, C., Isukapalli, S., Tang, X., Georgopoulos, P., He, S., Zalewsky, E., Hao, W., Ku, J.,
1103 Key, T., and Sistla, G. (2011), Impact of biogenic emission uncertainties on the simulated
1104 response of ozone and fine Particulate Matter to anthropogenic emission reductions, *J. Air*
1105 *Waste Manage.*, 61, 92–108.

1106 Huang, M., Carmichael, G. R., Adhikary, B., Spak, S. N., Kulkarni, S., Cheng, Y. F., Wei, C.,
1107 Tang, Y., Parrish, D. D., Oltmans, S. J., D'Allura, A., Kaduwela, A., Cai, C.,
1108 Weinheimer, A. J., Wong, M., Pierce, R. B., Al-Saadi, J. A., Streets, D. G., and Zhang, Q.
1109 (2010), Impacts of transported background ozone on California air quality during the
1110 ARCTAS-CARB period – a multi-scale modeling study, *Atmos. Chem. Phys.*, 10, 6947–6968,
1111 doi: 10.5194/acp-10-6947-2010.

1112 Huang, M., Carmichael, G. R., Chai, T., Pierce, R. B., Oltmans, S. J., Jaffe, D. A.,
1113 Bowman, K. W., Kaduwela, A., Cai, C., Spak, S. N., Weinheimer, A. J., Huey, L. G., and
1114 Diskin, G. S. (2013a), Impacts of transported background pollutants on summertime western
1115 US air quality: model evaluation, sensitivity analysis and data assimilation, *Atmos. Chem.*
1116 *Phys.*, 13, 359–391, doi: 10.5194/acp-13-359-2013.

1117 Huang, M., Bowman, K. W., Carmichael, G. R., Pierce, R. B., Worden, H. M., Luo, M., Cooper,
1118 O. R., Pollack, I. B., Ryerson, T. B., Brown, S. S. (2013b), Impact of southern California
1119 anthropogenic emissions on ozone pollution in the mountain states, *J. Geophys. Res.*, 118,
1120 12784-12803, doi: 10.1002/2013JD020205.

1121 Huang, M., et al. (2014), Changes in nitrogen oxides emissions in California during 2005–2010
1122 indicated from top-down and bottom-up emission estimates, *J. Geophys. Res.*, 119, 12,928–
1123 12,952, doi: 10.1002/2014JD022268, 2014.

1124 Huang, M., et al. (2015), Improved Western US Background Ozone Estimates via Constraining
1125 Nonlocal and Local Source Contributions using Aura TES and OMI Observations, *J. Geophys.*
1126 *Res.*, 120, 3572–3592, doi: 10.1002/2014JD022993.

1127 Huang, M., Carmichael, G. R., Crawford, J. H., Wisthaler, A., Zhan, X., Hain, C. R., Lee, P., and
1128 Guenther, A. B. (2017), Linkages between land initialization of the NASA-Unified WRF v7
1129 and biogenic isoprene emission estimates during the SEAC4RS and DISCOVER-AQ airborne
1130 campaigns, *Geosci. Model Dev. Discuss.*, doi:10.5194/gmd-2017-13, in review.

1131 Inness, A., Blechschmidt, A.-M., Bouarar, I., Chabrillat, S., Crepulja, M., Engelen, R. J., Eskes,
1132 H., Flemming, J., Gaudel, A., Hendrick, F., Huijnen, V., Jones, L., Kapsomenakis, J.,
1133 Katragkou, E., Keppens, A., Langerock, B., de Mazière, M., Melas, D., Parrington, M., Peuch,
1134 V. H., Razinger, M., Richter, A., Schultz, M. G., Suttie, M., Thouret, V., Vrekoussis, M.,
1135 Wagner, A., and Zerefos, C. (2015), Data assimilation of satellite-retrieved ozone, carbon
1136 monoxide and nitrogen dioxide with ECMWF's Composition-IFS, *Atmos. Chem. Phys.*, 15,
1137 5275-5303, doi:10.5194/acp-15-5275-2015.

1138 Jaffe, D.A. (2011), Relationship between surface and free tropospheric ozone in the Western U.S.,
1139 *Environ. Sci. Technol.*, 45, 432–438, doi: 10.1021/es1028102.

1140 Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Dentener, F., Muntean, M., Pouliot, G.,
1141 Keating, T., Zhang, Q., Kurokawa, J., Wankmüller, R., Denier van der Gon, H., Kuenen, J. J.
1142 P., Klimont, Z., Frost, G., Darras, S., Koffi, B., and Li, M. (2015), HTAP_v2.2: a mosaic of
1143 regional and global emission grid maps for 2008 and 2010 to study hemispheric transport of
1144 air pollution, *Atmos. Chem. Phys.*, 15, 11411-11432, doi:10.5194/acp-15-11411-2015.

1145 Jacob, D. J., Logan, J. A., and Murti, P. P. (1999), Effect of rising Asian emissions on surface
1146 ozone in the United States, *Geophys. Res. Lett.*, 26, 2175-2178, doi: 10.1029/1999GL900450.

1147 Jerret, M., R. T. Burnett, C. A. Popo, III, K. Ito, G. Thurston, D. Krewski, Y. Shi, E. Calle, and M.
1148 Thun (2009), Long-Term Ozone Exposure and Mortality, *the New England Journal of*
1149 *Medicine*, 360, 1085-1096, doi: 10.1056/NEJMoa0803894.

1150 Kaiser, J. W., Heil, A., Andreae, M. O., Benedetti, A., Chubarova, N., Jones, L., Morcrette, J.-J.,
1151 Razinger, M., Schultz, M. G., Suttie, M., and van der Werf, G. R. (2012), Biomass burning
1152 emissions estimated with a global fire assimilation system based on observed fire radiative
1153 power, *Biogeosciences*, 9, 527–554, doi:10.5194/bg-9-527-2012.

1154 Kalnay, E., and Co-authors (1996), The NCEP/NCAR 40-Year Reanalysis Project, *Bulletin of the*
1155 *American Meteorological Society*, 77, 437–471.

1156 Kim, S.-W., B. C. McDonald, S. Baidar, S. S. Brown, B. Dube, R. A. Ferrare, G. J. Frost, R. A.
1157 Harley, J. S. Holloway, H.-J. Lee, et al. (2016), Modeling the weekly cycle of NO_x and CO
1158 emissions and their impacts on O₃ in the Los Angeles-South Coast Air Basin during the CalNex
1159 2010 field campaign, *J. Geophys. Res. Atmos.*, 121, 1340–1360, doi:10.1002/2015JD024292.

1160 Koffi, B., Dentener, F., Janssens-Maenhout, G., Guizzardi, D., Crippa, M., Diehl, T., Galmarini,
1161 S., and Solazzo, E.: Hemispheric Transport Air Pollution (HTAP): Specification of the HTAP2

1162 experiments – Ensuring harmonized modelling, EUR 28255 EN – Scientific and Technical
1163 Research Reports, doi:10.2788/725244, 2016.

1164 Langford, A. O., Brioude, J., Cooper, O.R., Senff, C.J., Alvarez II, R.J., Hardesty, R.M., Johnson,
1165 B.J., and Oltmans, S.J. (2011), Stratospheric influence on surface ozone in the Los Angeles
1166 area during late spring and early summer of 2010, *J. Geophys. Res. Atmos.*, 117, D00V06, doi:
1167 10.1029/2011JD016766.

1168 Lapina, K., D. K. Henze, J. B. Milford, M. Huang, M. Lin, A. M. Fiore, G. Carmichael, G. G.
1169 Pfister, and K. Bowman (2014), Assessment of source contributions to seasonal vegetative
1170 exposure to ozone in the U.S., *J. Geophys. Res. Atmos.*, 119, 324–340,
1171 doi:10.1002/2013JD020905.

1172 Levelt, P.F., E. Hilsenrath, G.W. Leppelmeier, G.H.J. van den Oord, P.K. Bhartia, J. Tamminen,
1173 J.F. de Haan and J.P. Veefkind (2006), Science Objectives of the Ozone Monitoring Instrument,
1174 *IEEE Transaction on Geoscience and Remote Sensing*, 44, 1199-1208.

1175 Li, M., Zhang, Q., Kurokawa, J.-I., Woo, J.-H., He, K., Lu, Z., Ohara, T., Song, Y., Streets, D. G.,
1176 Carmichael, G. R., Cheng, Y., Hong, C., Huo, H., Jiang, X., Kang, S., Liu, F., Su, H., and
1177 Zheng, B. (2017), MIX: a mosaic Asian anthropogenic emission inventory under the
1178 international collaboration framework of the MICS-Asia and HTAP, *Atmos. Chem. Phys.*, 17,
1179 935-963, doi:10.5194/acp-17-935-2017.

1180 Lin, M., Holloway, T., Carmichael, G. R., and Fiore, A. M. (2010), Quantifying pollution inflow
1181 and outflow over East Asia in spring with regional and global models, *Atmos. Chem. Phys.*,
1182 10, 4221-4239, doi:10.5194/acp-10-4221-2010.

1183 Lin, M., A. M. Fiore, L. W. Horowitz, O. R. Cooper, V. Naik, J. Holloway, B. J. Johnson, A.
1184 Middlebrook, S. J. Oltmans, I. B. Pollack, T. B. Ryerson, J. X. Warner, C. Wiedinmyer, J.
1185 Wilson, B. Wyman (2012a), Transport of Asian ozone pollution into surface air over the
1186 western United States in spring, *J. Geophys. Res.*, 117, D00V07, doi: 10.1029/2011JD016961.

1187 Lin, M., A. Fiore, O. R. R. Cooper, L. W. Horowitz, A. O. O. Langford, H. Levy II, B. J. Johnson,
1188 V. Naik, S. J. Oltmans, and C. J. Senff (2012b), Springtime high surface ozone events over the
1189 western United States: Quantifying the role of stratospheric intrusions, *J. Geophys. Res.*, 117,
1190 D00V22, doi: 10.1029/2012JD018151.

1191 Lin, M., L.W. Horowitz, S. J. Oltmans, A. M. Fiore, S. Fan (2014), Tropospheric ozone trends at
1192 Manna Loa Observatory tied to decadal climate variability, *Nature Geoscience*, 7, 136-143,
1193 doi:10.1038/NGEO2066.

1194 Lin, M., L. W. Horowitz, O. R. Cooper, D. Tarasick, S. Conley, L. T. Iraci, B. Johnson, T. Leblanc,
1195 I. Petropavlovskikh, and E. L. Yates (2015), Revisiting the evidence of increasing springtime
1196 ozone mixing ratios in the free troposphere over western North America, *Geophys. Res. Lett.*,
1197 42, 8719–8728, doi:10.1002/2015GL065311.

1198 Lin, M., Horowitz, L. W., Payton, R., Fiore, A. M., and Tonnesen, G. (2016), US surface ozone
1199 trends and extremes from 1980–2014: Quantifying the roles of rising Asian emissions,
1200 domestic controls, wildfires, and climate, *Atmos. Chem. Phys. Discuss.*, doi:10.5194/acp-
1201 2016-1093, in review.

1202 Liu, F., Q. Zhang, R. J. van der A, B. Zheng, D. Tong, L. Yan, Y. Zheng, and K. He (2016), Recent
1203 reduction in NO_x emissions over China: Synthesis of satellite observations and emission
1204 inventories, *Environ. Res. Lett.*, 11 (11), 114002, doi: 10.1088/1748-9326/11/11/114002.

1205 Livesey, N.J., M.J. Filipiak, L. Froidevaux, W.G. Read, A. Lambert, M.L. Santee, J.H. Jiang, H.C.
1206 Pumphrey, J.W. Waters, R.E. Cofield, D.T. Cuddy, W.H. Daffer, B.J. Drouin, R.A. Fuller, R.F.
1207 Jarnot, Y.B. Jiang, B.W. Knosp, Q.B. Li, V.S. Perun, M.J. Schwartz, W.V. Snyder, P.C. Stek,

1208 R.P. Thurstans, P.A. Wagner, M. Avery, E.V. Browell, J-P. Cammas, L.E. Christensen, G.S.
1209 Diskin, R-S. Gao, H-J. Jost, M. Loewenstein, J.D. Lopez, P. Nedelec, G.B. Osterman, G.W.
1210 Sachse, and C.R. Webster (2008), Validation of Aura Microwave Limb Sounder O₃ and CO
1211 observations in the upper troposphere and lower stratosphere, *J. Geophys. Res.* 113, D15S02,
1212 doi:10.1029/2007JD008805.

1213 Luecken, D.J., S. Phillips, G. Sarwar, C. Jang, Effects of using the CB05 vs. SAPRC99 vs. CB4
1214 chemical mechanism on model predictions (2008), Ozone and gas-phase photochemical
1215 precursor concentrations, *Atmos. Environ.*, 42 (23), 5805-5820, doi:
1216 10.1016/j.atmosenv.2007.08.056.

1217 Maas, R. and P. Grennfelt (eds) (2016), Towards Cleaner Air Scientific Assessment Report 2016.
1218 EMEP Steering Body and Working Group on Effects of the Convention on Long-Range
1219 Transboundary Air Pollution, Oslo,
1220 http://www.unece.org/fileadmin/DAM/env/Irtap/ExecutiveBody/35th_session/CLRTAP_Scientific_Assessment_Report_-_Final_20-5-2016.pdf.

1221 Madronich, S., Flocke, S., Zeng, J., Petropavlovskikh, I., and Lee-Taylor, J. (2002), The
1222 Tropospheric Ultra-violet Visible (TUV) model Manual,
1223 [https://www2.acom.ucar.edu/modeling/tropospheric-ultraviolet-and-visible-tuv-radiation-](https://www2.acom.ucar.edu/modeling/tropospheric-ultraviolet-and-visible-tuv-radiation-model)
1224 [model](https://www2.acom.ucar.edu/modeling/tropospheric-ultraviolet-and-visible-tuv-radiation-model).

1226 Mauzerall, D. L. and Wang, X. (2001), Protecting Agricultural Crops from the Effects of
1227 Tropospheric Ozone Exposure: Reconciling Science and Standard Setting in the United States,
1228 Europe and Asia, *Annual Review of Energy and the Environment*, 26, 237-268.

1229 McDonald-Buller, E. C., et al. (2011), Establishing policy relevant background (PRB) ozone
1230 concentrations in the United States, *Environ. Sci. Technol.*, 45, 9484–9497.

1231 Meijer, E. W., van Velthoven, P. F. J., Brunner, D. W., Huntrieser, H., and Kelder, H. (2001),
1232 Improvement and evaluation of the parameterization of nitrogen oxide production by lightning,
1233 *Phys. Chem. Earth Pt. C*, 26, 577–583.

1234 Mesinger, F., DiMego, G., Kalnay, E., Mitchell, K., Shafran, P. C., Ebisuzaki, W., Jovic, D.,
1235 Woollen, J., Rogers, E., Berbery, E. H., Ek, M. B., Fan, Y., Grumbine, R., Higgins, W., Li, H.,
1236 Lin, Y., Manikin, G., Parrish, D. and Shi, W. (2006), North American Regional Reanalysis,
1237 *Bulletin of the American Meteorological Society*, 87(3), 343–360, doi: 10.1175/BAMS-87-3-
1238 343.

1239 Miyazaki, K., Eskes, H. J., Sudo, K., Takigawa, M., van Weele, M., Boersma, K. F. (2012),
1240 Simultaneous assimilation of satellite NO₂, O₃, CO, and HNO₃ data for the analysis of
1241 tropospheric chemical composition and emissions, *Atmos. Chem. Phys.*, 12, 9545-9579.

1242 Monks, P. S., Archibald, A. T., Colette, A., Cooper, O., Coyle, M., Derwent, R., Fowler, D.,
1243 Granier, C., Law, K. S., Mills, G. E., Stevenson, D. S., Tarasova, O., Thouret, V., von
1244 Schneidmesser, E., Sommariva, R., Wild, O., and Williams, M. L. (2015), Tropospheric
1245 ozone and its precursors from the urban to the global scale from air quality to short-lived
1246 climate forcer, *Atmos. Chem. Phys.*, 15, 8889-8973, doi:10.5194/acp-15-8889-2015.

1247 Murray, L. T., D. J. Jacob, J. A. Logan, R. C. Hudman, and W. J. Koshak (2012), Optimized
1248 regional and interannual variability of lightning in a global chemical transport model
1249 constrained by LIS/OTD satellite data, *J. Geophys. Res.*, 117, D20307,
1250 doi:10.1029/2012JD017934.

1251 National Research Council (NRC) (2009), global sources of local pollution-An Assessment of
1252 Long-Range Transport of Key Air Pollutants to and from the United States, 35-66,
1253 http://books.nap.edu/openbook.php?record_id=12743&page=35.

1254 Neuman, J. A., et al. (2012), Observations of ozone transport from the free troposphere to the Los
1255 Angeles basin, *J. Geophys. Res. Atmos.*, 117, D00V09, doi: 10.1029/2011JD016919.

1256 Oetjen, H., Payne, V. H., Kulawik, S. S., Eldering, A., Worden, J., Edwards, D. P., Francis, G. L.,
1257 Worden, H. M., Clerbaux, C., Hadji-Lazaro, J., and Hurtmans, D. (2014), Extending the
1258 satellite data record of tropospheric ozone profiles from Aura-TES to MetOp-IASI:
1259 characterisation of optimal estimation retrievals, *Atmos. Meas. Tech.*, 7, 4223–4236,
1260 doi:10.5194/amt-7-4223-2014.

1261 Oetjen, H., Payne, V. H., Neu, J. L., Kulawik, S. S., Edwards, D. P., Eldering, A., Worden, H. M.,
1262 and Worden, J. R. (2016), A joint data record of tropospheric ozone from Aura-TES and
1263 MetOp-IASI, *Atmos. Chem. Phys.*, 16, 10229–10239, doi:10.5194/acp-16-10229-2016.

1264 Ott, L. E., B. N. Duncan, A. M. Thompson, G. Diskin, Z. Fasnacht, A. O. Langford, M. Lin, A. M.
1265 Molod, J. E. Nielsen, S. E. Pusede, et al. (2016), Frequency and impact of summertime
1266 stratospheric intrusions over Maryland during DISCOVER-AQ (2011): New evidence from
1267 NASA's GEOS-5 simulations, *J. Geophys. Res. Atmos.*, 121, 3687–3706,
1268 doi:10.1002/2015JD024052.

1269 Park, R. J., D. J. Jacob, B. D. Field, R. M. Yantosca, and M. Chin (2004), Natural and
1270 transboundary pollution influences on sulfate-nitrate-ammonium aerosols in the United States:
1271 Implications for policy, *J. Geophys. Res.*, 109, D15204, doi:10.1029/2003JD004473.

1272 Parrington, M., D. B. A. Jones, K. W. Bowman, L. W. Horowitz, A. M. Thompson, D. W. Tarasick,
1273 and J. C. Witte (2008), Estimating the summertime tropospheric ozone distribution over North
1274 America through assimilation of observations from the Tropospheric Emission Spectrometer,
1275 *J. Geophys. Res.*, 113, D18307, doi:10.1029/2007JD009341.

1276 Parrington, M., D. B. A. Jones, K. W. Bowman, A. M. Thompson, D. W. Tarasick, J. Merrill, S.
1277 J. Oltmans, T. Leblanc, J. C. Witte, and D. B. Millet (2009), Impact of the assimilation of
1278 ozone from the Tropospheric Emission Spectrometer on surface ozone across North America,
1279 *Geophys. Res. Lett.*, 36, L04802, doi:10.1029/2008GL036935.

1280 Parrish, D. D., D. B. Millet, and A. H. Goldstein (2009), Increasing ozone in marine boundary
1281 layer inflow at the west coasts of North America and Europe, *Atmos. Chem. Phys.*, 9, 1303–
1282 1323, doi:10.5194/acp-9-1303-2009.

1283 Parrish, D. D., Aikin, K. C., Oltmans, S. J., Johnson, B. J., Ives, M., and Sweeny, C. (2010), Impact
1284 of transported background ozone inflow on summertime air quality in a California ozone
1285 exceedance area, *Atmos. Chem. Phys.*, 10, 10093–10109, doi:10.5194/acp-10-10093-2010.

1286 Parrish, D. D., et al. (2012), Long-term changes in lower tropospheric baseline ozone
1287 concentrations at northern mid-latitudes, *Atmos. Chem. Phys.*, 12, 11,485–11,504,
1288 doi:10.5194/acp-12-11485-2012.

1289 Pierce, R. B., et al. (2007), Chemical data assimilation estimates of continental U.S. ozone and
1290 nitrogen budgets during the Intercontinental Chemical Transport Experiment–North America,
1291 *J. Geophys. Res.*, 112, D12S21, doi:10.1029/2006JD007722.

1292 Pierce, R. B., et al. (2009), Impacts of background ozone production on Houston and Dallas, Texas,
1293 air quality during the Second Texas Air Quality Study field mission, *J. Geophys. Res.*, 114,
1294 D00F09, doi:10.1029/2008JD011337.

1295 Pouliot, G., H. A.C. Denier van der Gon, J. Kuenen, J. Zhang, M. D. Moran, P.A. Makar (2015),
1296 Analysis of the emission inventories and model-ready emission datasets of Europe and North
1297 America for phase 2 of the AQMEII project, *Atmos. Environ.*, 115, 345–360.

1298 Qu, Z., D. K. Henze, S. L. Capps, Y. Wang, X. Xu, J. Wang (2016), Monthly top-down NO_x
1299 emissions for China (2005–2012): a hybrid inversion method and trend analysis, submitted.

1300 Quennehen, B., Raut, J.-C., Law, K. S., Daskalakis, N., Ancellet, G., Clerbaux, C., Kim, S.-W.,
 1301 Lund, M. T., Myhre, G., Olivié, D. J. L., Safieddine, S., Skeie, R. B., Thomas, J. L., Tsyro, S.,
 1302 Bazureau, A., Bellouin, N., Hu, M., Kanakidou, M., Klimont, Z., Kupiainen, K.,
 1303 Myriokefalitakis, S., Quaas, J., Rumbold, S. T., Schulz, M., Cherian, R., Shimizu, A., Wang,
 1304 J., Yoon, S.-C., and Zhu, T. (2016), Multi-model evaluation of short-lived pollutant
 1305 distributions over east Asia during summer 2008, *Atmos. Chem. Phys.*, 16, 10765-10792,
 1306 doi:10.5194/acp-16-10765-2016.

1307 Reidmiller, D. R., Fiore, A. M., Jaffe, D. A., Bergmann, D., Cuvelier, C., Dentener, F. J., Duncan,
 1308 B. N., Folberth, G., Gauss, M., Gong, S., Hess, P., Jonson, J. E., Keating, T., Lupu, A., Marmer,
 1309 E., Park, R., Schultz, M. G., Shindell, D. T., Szopa, S., Vivanco, M. G., Wild, O., and Zuber,
 1310 A. (2009), The influence of foreign vs. North American emissions on surface ozone in the US,
 1311 *Atmos. Chem. Phys.*, 9, 5027-5042, doi:10.5194/acp-9-5027-2009.

1312 Rodgers, C. D. (2000), *Inverse Methods for Atmospheric Sounding: Theory and Practice*, World
 1313 Sci., Singapore.

1314 Ryerson, T. B., Andrews, A. E., Angevine, W. M., Bates, T. S., Brock, C. A., Cairns, B., Cohen,
 1315 R. C., Cooper, O. R., de Gouw, J. A., Fehsenfeld, F. C., Ferrare, R. A., Fischer, M. L., Flagan,
 1316 R. C., Goldstein, A. H., Hair, J. W., Hardesty, R. M., Hostetler, C. A., Jimenez, J. L., Langford,
 1317 A. O., McCauley, E., McKeen, S. A., Molina, L. T., Nenes, A., Oltmans, S. J., Parrish, D. D.,
 1318 Pederson, J. R., Pierce, R. B., Prather, K., Quinn, P. K., Seinfeld, J. H., Senff, C. J., Sorooshian,
 1319 A., Stutz, J., Surratt, J. D., Trainer, M., Volkamer, R., Williams, E. J., Wofsy, S. C. (2013),
 1320 The 2010 California Research at the Nexus of Air Quality and Climate Change (CalNex) field
 1321 study, *J. Geophys. Res.*, 118, 5830–5866.

1322 Schere, K. J., Flemming, R., Vautard, C., Chemel, A., Colette, C., Hogrefe, B., Bessagnet, F., Meleux,
 1323 R. Mathur, S. Roselle, R.-M. Hu, R. S. Sokhi, S. T. Rao, S. Galmarini (2012), Trace gas/aerosol
 1324 boundary concentrations and their impacts on continental-scale AQMEII modeling domains,
 1325 *Atmos. Environ.*, 53, 38-50, doi: 10.1016/j.atmosenv.2011.09.043.

1326 Shindell, D. T., G. Faluvegi, D. M. Koch, G. A. Schmidt, N. Unger, and S. E. Bauer (2009),
 1327 Improved attribution of climate forcing to emissions, *Science*, 326, 716–718, doi:
 1328 10.1126/science.1174760.

1329 Shindell, D. T., et al. (2013), Radiative forcing in the ACCMIP historical and future climate
 1330 simulations, *Atmos. Chem. Phys.*, 13, 2939–2974, doi:10.5194/acp-13-2939-2013.

1331 Simpson, D., Benedictow, A., Berge, H., Bergström, R., Emberson, L. D., Fagerli, H., Flechard,
 1332 C. R., Hayman, G. D., Gauss, M., Jonson, J. E., Jenkin, M. E., Nyíri, A., Richter, C., Semeena,
 1333 V. S., Tsyro, S., Tuovinen, J.-P., Valdebenito, Á., and Wind, P. (2012), The EMEP MSC-W
 1334 chemical transport model – technical description, *Atmos. Chem. Phys.*, 12, 7825–7865,
 1335 doi:10.5194/acp-12-7825-2012.

1336 Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrou, T., Müller, J.-F.,
 1337 Kuhn, U., Stefani, P., and Knorr, W. (2014), Global data set of biogenic VOC emissions
 1338 calculated by the MEGAN model over the last 30 years, *Atmos. Chem. Phys.*, 14, 9317–9341,
 1339 doi:10.5194/acp-14-9317-2014.

1340 Skamarock, W. C., J. B. Klemp, J. Dudhia, D. Gill, D. M. Barker, W. Wang, and J. G. Powers
 1341 (2008), A description of the Advanced Research WRF version 3 (Available at
 1342 www.mmm.ucar.edu/wrf/users/docs/arwv3.pdf).

1343 Smith, K. R., Jerrett, M., and Anderson, H. R. et al. (2009), Public health benefits of strategies to
 1344 reduce greenhouse-gas emissions: health implications of short-lived greenhouse pollutants,
 1345 *Lancet*, doi: 10.1016/S0140-6736 (09) 61716-5.

1346 Solazzo, E. R. Bianconi, R. Vautard, K. W. Appel, M. D. Moran, C. Hogrefe, B. Bessagnet, J.
1347 Brandt, J. H. Christensen, C. Chemel, I. Coll, H. D. van der Gon, J. Ferreira, R. Forkel, X. V.
1348 Francis, G. Grell, P. Grossi, A. B. Hansen, A. Jeričević, L. Kraljević, A. I. Miranda, U.
1349 Nopmongcol, G. Pirovano, M. Prank, A. Riccio, K. N. Sartelet, M. Schaap, J. D. Silver, R. S.
1350 Sokhi, J. Vira, J. Werhahn, R. Wolke, G. Yarwood, J. Zhang, S.T. Rao, S. Galmarini (2012),
1351 Model evaluation and ensemble modelling of surface-level ozone in Europe and North
1352 America in the context of AQMEII, *Atmos. Environ.*, 53, 60-74, , doi:
1353 10.1016/j.atmosenv.2012.01.003.

1354 Søvde, O. A., Prather, M. J., Isaksen, I. S. A., Berntsen, T. K., Stordal, F., Zhu, X., Holmes, C. D.,
1355 and Hsu, J. (2012), The chemical transport model Oslo CTM3, *Geosci. Model Dev.*, 5, 1441–
1356 1469, doi:10.5194/gmd-5-1441-2012.

1357 Sudo, K., M. Takahashi, J. Kurokawa, and H. Akimoto (2002), Chaser: A global chemical model
1358 of the troposphere, 1. Model description, *J. Geophys. Res.*, 107(D17), 4339,
1359 doi:10.1029/2001JD001113.

1360 Stevenson, D. S., et al. (2006), Multimodel ensemble simulations of present-day and near-future
1361 tropospheric ozone, *J. Geophys. Res.*, 111, D08301, doi:10.1029/2005JD006338.

1362 Stevenson, D. S., et al. (2013), Tropospheric ozone changes, radiative forcing and attribution to
1363 emissions in the Atmospheric Chemistry and Climate Model Intercomparison Project
1364 (ACCMIP), *Atmos. Chem. Phys.*, 13, 3063–3085, doi:10.5194/acp-13-3063-2013.

1365 Susaya, J., Kim, K.-H., Shon, Z.-H., Brown R. J. (2013), Demonstration of long-term increases in
1366 tropospheric O₃ levels: Causes and potential impacts, *Chemosphere*, 92, 1520–1528.

1367 Task Force on Hemispheric Transport of Air Pollution (HTAP) (2010), 2010 Final Assessment
1368 report, Part A: Ozone and particulate matter,
1369 [http://www.htap.org/activities/2010_Final_Report/HTAP%202010%20Part%20A%2011040](http://www.htap.org/activities/2010_Final_Report/HTAP%202010%20Part%20A%20110407.pdf)
1370 [7.pdf](http://www.htap.org/activities/2010_Final_Report/HTAP%202010%20Part%20A%20110407.pdf).

1371 Tilmes, S., Lamarque, J.-F., Emmons, L. K., Kinnison, D. E., Marsh, D., Garcia, R. R., Smith, A.
1372 K., Neely, R. R., Conley, A., Vitt, F., Val Martin, M., Tanimoto, H., Simpson, I., Blake, D. R.,
1373 and Blake, N. (2016), Representation of the Community Earth System Model (CESM1)
1374 CAM4-chem within the Chemistry- Climate Model Initiative (CCMI), *Geosci. Model Dev.*, 9,
1375 1853– 1890, doi:10.5194/gmd-9-1853-2016.

1376 Travis, K. R., Jacob, D. J., Fisher, J. A., Kim, P. S., Marais, E. A., Zhu, L., Yu, K., Miller, C. C.,
1377 Yantosca, R. M., Sulprizio, M. P., Thompson, A. M., Wennberg, P. O., Crounse, J. D., St.
1378 Clair, J. M., Cohen, R. C., Laughner, J. L., Dibb, J. E., Hall, S. R., Ullmann, K., Wolfe, G. M.,
1379 Pollack, I. B., Peischl, J., Neuman, J. A., and Zhou, X. (2016), Why do models overestimate
1380 surface ozone in the Southeast United States?, *Atmos. Chem. Phys.*, 16, 13561-13577,
1381 doi:10.5194/acp-16-13561-2016.

1382 United Nations Environment Programme and World Meteorological Organization (2011),
1383 Integrated Assessment of Black Carbon and Tropospheric Ozone: Summary for Decision
1384 Makers, http://www.unep.org/dewa/Portals/67/pdf/Black_Carbon.pdf.

1385 US EPA (2016a), Implementation of the 2015 Primary Ozone NAAQS: Issues Associated with
1386 Background Ozone White Paper for Discussion,
1387 <https://www.epa.gov/sites/production/files/2016-03/documents/whitepaper-bgo3-final.pdf>.

1388 US EPA (2016b), High level summary of background ozone workshop,
1389 [https://www.epa.gov/sites/production/files/2016-03/documents/bgo3-high-level-](https://www.epa.gov/sites/production/files/2016-03/documents/bgo3-high-level-summary.pdf)
1390 [summary.pdf](https://www.epa.gov/sites/production/files/2016-03/documents/bgo3-high-level-summary.pdf).

1391 van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton,
1392 D. C., DeFries, R. S., Jin, Y., and van Leeuwen, T. T. (2010), Global fire emissions and the
1393 contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), *Atmos.*
1394 *Chem. Phys.*, 10, 11707–11735, doi:10.5194/acp-10-11707-2010.

1395 van Noije, T. P. C., Eskes, H. J., Dentener, F. J., Stevenson, D. S., Ellingsen, K., Schultz, M. G.,
1396 Wild, O., Amann, M., Atherton, C. S., Bergmann, D. J., Bey, I., Boersma, K. F., Butler, T.,
1397 Cofala, J., Drevet, J., Fiore, A. M., Gauss, M., Hauglustaine, D. A., Horowitz, L. W., Isaksen,
1398 I. S. A., Krol, M. C., Lamarque, J.-F., Lawrence, M. G., Martin, R. V., Montanaro, V., Müller,
1399 J.-F., Pitari, G., Prather, M. J., Pyle, J. A., Richter, A., Rodriguez, J. M., Savage, N. H., Strahan,
1400 S. E., Sudo, K., Szopa, S., and van Roozendaal, M. (2006), Multi-model ensemble simulations
1401 of tropospheric NO₂ compared with GOME retrievals for the year 2000, *Atmos. Chem. Phys.*,
1402 6, 2943–2979, doi:10.5194/acp-6-2943-2006.

1403 Verstraeten, W. W., K. F. Boersma, J. Zörner, M. A. F. Allaart, K. W. Bowman, and J. R. Worden
1404 (2013), Validation of six years of TES tropospheric ozone retrievals with ozonesonde
1405 measurements: Implications for spatial patterns and temporal stability in the bias, *Atmos. Meas.*
1406 *Tech.*, 6, 1413–1423.

1407 Verstraeten, W. W., J. L. Neu, J. E. Williams, K. W. Bowman, J. R. Worden, and K. F. Boersma
1408 (2015), Rapid increases in tropospheric ozone production and export from China, *Nature*
1409 *Geoscience*, 8, 690–695, doi:10.1038/ngeo2493.

1410 Wang, H., D. J. Jacob, P. L. Sager, D. G. Streets, R. J. Park, A. B. Gilliland, and A. van Donkelaar
1411 (2009), Surface ozone background in the United States: Canadian and Mexican pollution
1412 influences, *Atmos. Environ.*, 43(6), 1310–1319, doi:10.1016/j.atmosenv.2008.11.036.

1413 Wang, Y., Konopka, P., Liu, Y., Chen, H., Müller, R., Plöger, F., Riese, M., Cai, Z., and Lü, D.
1414 (2012), Tropospheric ozone trend over Beijing from 2002–2010: ozonesonde measurements
1415 and modeling analysis, *Atmos. Chem. Phys.*, 12, 8389–8399, doi:10.5194/acp-12-8389-2012.

1416 Warneke, C., J. A. deGouw, J. S. Holloway, J. Peischl, T. B. Ryerson, E. Atlas, D. Blake, M.
1417 Trainer, and D. D. Parrish (2012), Multiyear trends in volatile organic compounds in Los
1418 Angeles, California: Five decades of decreasing emissions, *J. Geophys. Res.*, 117, D00V17,
1419 doi:10.1029/2012JD017899.

1420 Warner, J. X., McCourt Comer, M., Barnet, C. D., McMillan, W. W., Wolf, W., Maddy, E., and
1421 Sachse, G. (2007), A comparison of satellite tropospheric carbon monoxide measurements
1422 from AIRS and MOPITT during INTEX-A, *J. Geophys. Res.*, 112, D12S17,
1423 doi:10.1029/2006JD007925, 2007.

1424 Wiedinmyer, C., Akagi, S. K., Yokelson, R. J., Emmons, L. K., Al-Saadi, J. A., Orlando, J. J., and
1425 Soja, A. J. (2011), The Fire INventory from NCAR (FINN): a high resolution global model to
1426 estimate the emissions from open burning, *Geosci. Model Dev.*, 4, 625–641, doi:10.5194/gmd-
1427 4-625-2011.

1428 Wigder, N.L., Jaffe, D.A., Herron-Thorpe, F.L., and Vaughan, J.K. (2013), Influence of daily
1429 variations in baseline ozone on urban air quality in the United States Pacific Northwest, *J.*
1430 *Geophys. Res.*, 118, 3343–3354, doi: 10.1029/2012JD018738.

1431 Wild, O., Fiore, A. M., Shindell, D. T., Doherty, R. M., Collins, W. J., Dentener, F. J., Schultz, M.
1432 G., Gong, S., MacKenzie, I. A., Zeng, G., Hess, P., Duncan, B. N., Bergmann, D. J., Szopa,
1433 S., Jonson, J. E., Keating, T. J., and Zuber, A. (2012), Modelling future changes in surface
1434 ozone: a parameterized approach, *Atmos. Chem. Phys.*, 12, 2037–2054, doi:10.5194/acp-12-
1435 2037-2012.

1436 Wu, S., B. N. Duncan, D. J. Jacob, A. M. Fiore, and O. Wild (2009), Chemical nonlinearities in
1437 relating intercontinental ozone pollution to anthropogenic emissions, *Geophys. Res. Lett.*, 36,
1438 L05806, doi:10.1029/2008GL036607.

1439 Yarwood, G., Rao, S., Yocke, M., and Whitten, G. (2005), Updates to the carbon bond chemical
1440 mechanism: CB05. Final report to the US EPA, EPA Report Number: RT-0400675.

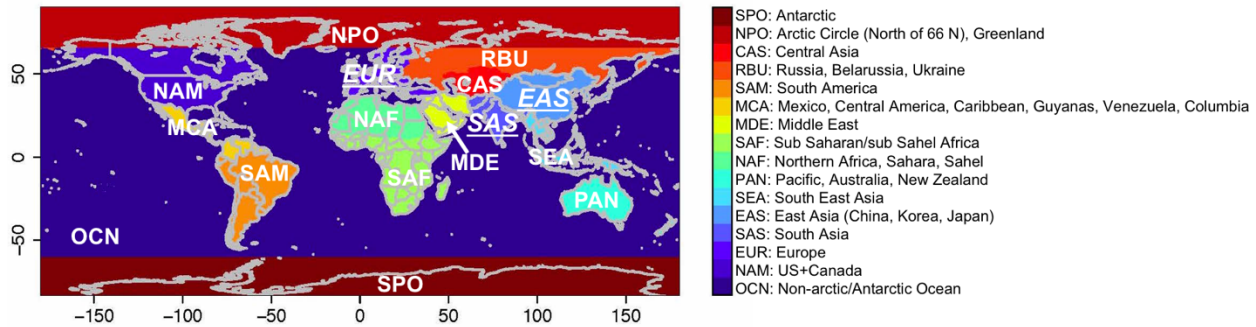
1441 Zhang, L., Jacob, D. J., Boersma, K. F., Jaffe, D. A., Olson, J. R., Bowman, K. W., Worden, J. R.,
1442 Thompson, A. M., Avery, M. A., Cohen, R. C., Dibb, J. E., Flock, F. M., Fuelberg, H. E.,
1443 Huey, L. G., McMillan, W. W., Singh, H. B., and Weinheimer, A. J. (2008), Transpacific
1444 transport of ozone pollution and the effect of recent Asian emission increases on air quality in
1445 North America: an integrated analysis using satellite, aircraft, ozonesonde, and surface
1446 observations, *Atmos. Chem. Phys.*, 8, 6117-6136, doi:10.5194/acp-8-6117-2008.

1447 Zhang, L., Jacob, D. J., Kopacz, M., Henze, D. K., Singh, K., and Jaffe, D. A. (2009),
1448 Intercontinental source attribution of ozone pollution at western U.S. sites using an adjoint
1449 method, *Geophys. Res. Lett.*, 36, L11810, doi: 10.1029/2009GL037950.

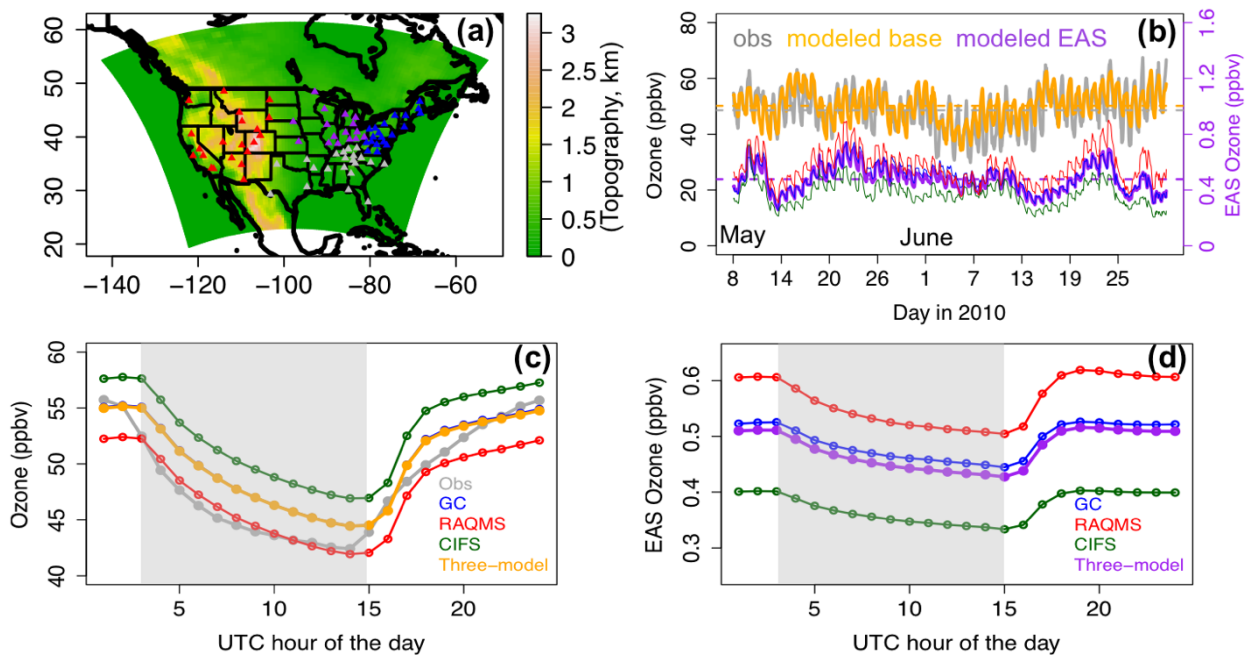
1450 Zhang, L., D. J. Jacob, N. V. Downey, D. A. Wood, D. Blewitt, C. C. Carouge, A. van Donkelaar,
1451 D. B. A. Jones, L. T. Murray, and Y. Wang (2011), Improved estimate of the policy-relevant
1452 background ozone in the United States using the GEOS-Chem global model with $1/2^{\circ} \times 2/3^{\circ}$
1453 horizontal resolution over North America, *Atmos. Environ.*, 45, 6769–6776, doi:
1454 10.1016/j.atmosenv.2011.07.054.

1455 Zhang, Q., Yuan, B., Shao, M., Wang, X., Lu, S., Lu, K., Wang, M., Chen, L., Chang, C.-C., and
1456 Liu, S. C. (2014), Variations of ground-level O₃ and its precursors in Beijing in summertime
1457 between 2005 and 2011, *Atmos. Chem. Phys.*, 14, 6089-6101, doi:10.5194/acp-14-6089-2014.

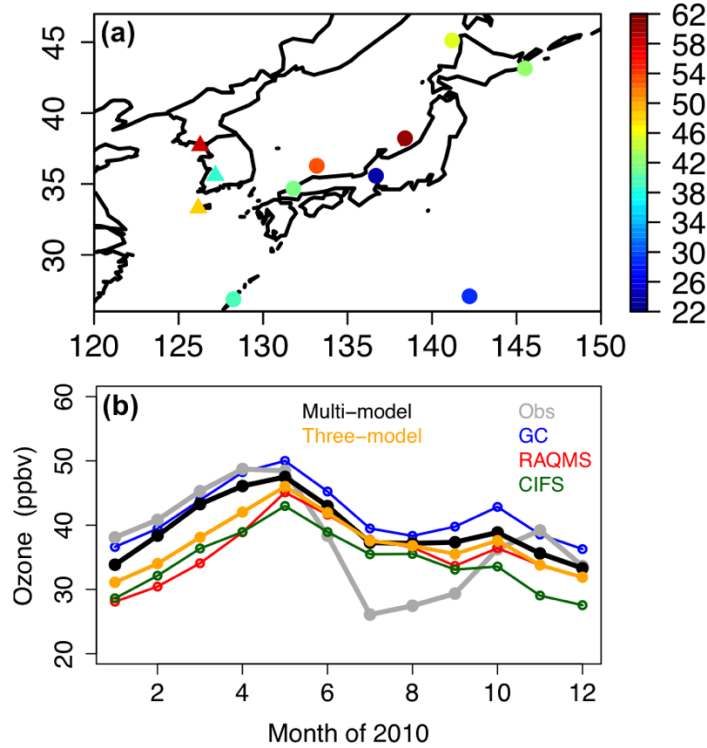
1458 Zhang, Y., Y. Chen, G. Sarwar, and K. Schere (2012), Impact of gas-phase mechanisms on
1459 Weather Research Forecasting Model with Chemistry (WRF/Chem) predictions: Mechanism
1460 implementation and comparative evaluation, *J. Geophys. Res.*, 117, D01301,
1461 doi:10.1029/2011JD015775.



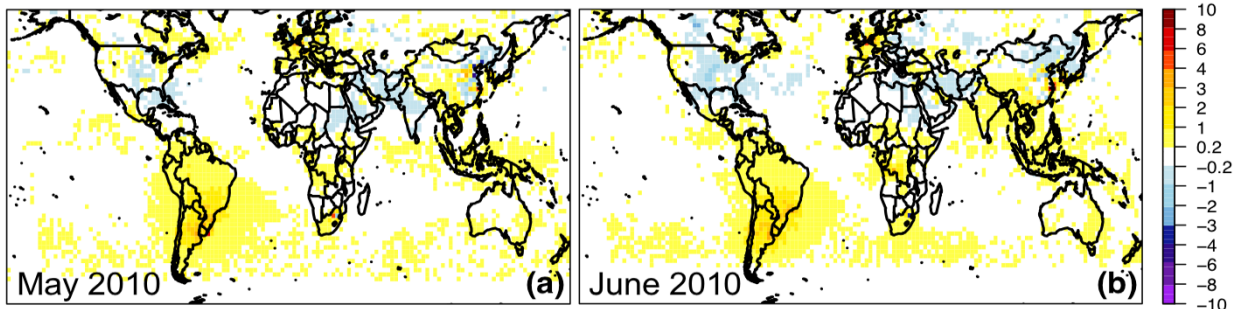
1462
 1463 **Figure 1.** Definitions of the 16 source regions used in HTAP2 SR relationship study (More details
 1464 in Koffi et al., 2016). The map is plotted based on data on a $0.1^\circ \times 0.1^\circ$ resolution grid. We focus
 1465 in this study on the impact of anthropogenic pollution from selected non-North American source
 1466 regions (i.e., EAS, SAS, and EUR), whose names are underlined and in italic.
 1467



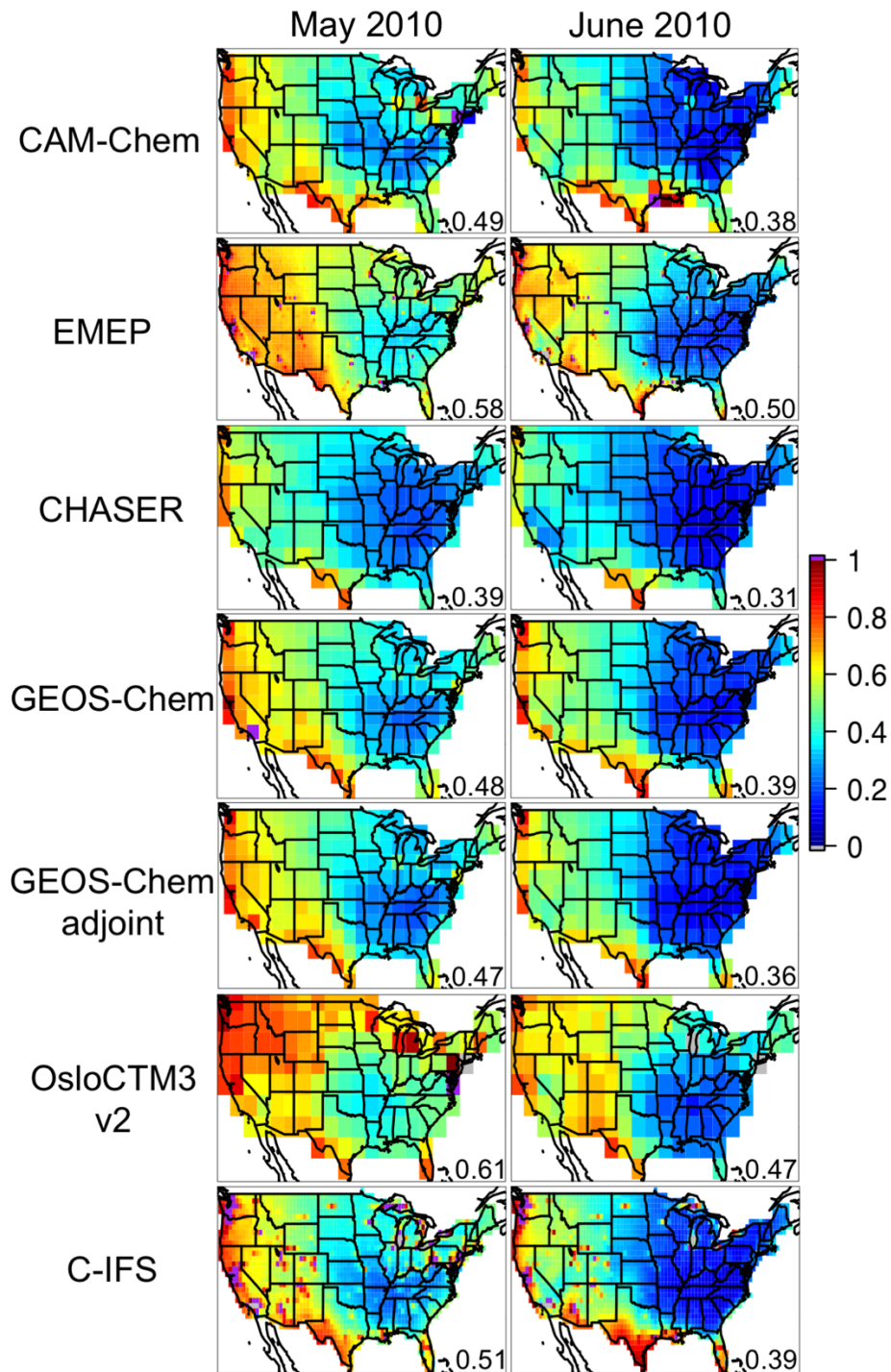
1468
 1469 **Figure 2.** (a) The 60 km STEM NAM domain, colored by the model topography. The CASTNET
 1470 sites used in the STEM base O₃ evaluation are marked as triangles in different colors that identify
 1471 the subregions they belong to (red: western US; grey: southern US; purple: Midwest; blue:
 1472 northeastern US). (b) Evaluation of the STEM modeled (averaged from the three base simulations
 1473 using the GEOS-Chem, ECMWF C-IFS, and RAQMS base runs as the chemical boundary
 1474 conditions) hourly O₃ at the western US (i.e., EPA regions 8, 9, and 10) CASTNET sites.
 1475 Observations, modeled base O₃ and the modeled R(O₃, EAS, 20%) are in grey, orange, and purple
 1476 lines, respectively. The horizontal dashed lines indicate the period mean values. The R(O₃, EAS,
 1477 20%) values from STEM calculations using three different chemical boundary conditions
 1478 are shown separately in thin lines (blue: GEOS-Chem; red: RAQMS; green: C-IFS). The period-mean
 1479 diurnal variability of the STEM modeled (c) base and (d) R(O₃, EAS, 20%) at the western US
 1480 CASTNET sites. The STEM calculations using three different chemical boundary conditions are
 1481 shown separately as well as averagely. Light grey-shaded areas indicate the local standard
 1482 nighttime (from 6/7 pm to 7/8 am).



1483
 1484 **Figure 3.** (a) May-June 2010 period mean surface O₃ observations in ppbv at eight Japanese (filled
 1485 circles) and three Korean (filled triangles) EANET sites. (b) Observed and modeled monthly-mean
 1486 surface O₃ in 2010 at all eleven EANET sites. The “Multi-model” and “Three-model” in the legend
 1487 indicate the mean values of all eight global models and only of the three boundary condition
 1488 models, respectively.
 1489

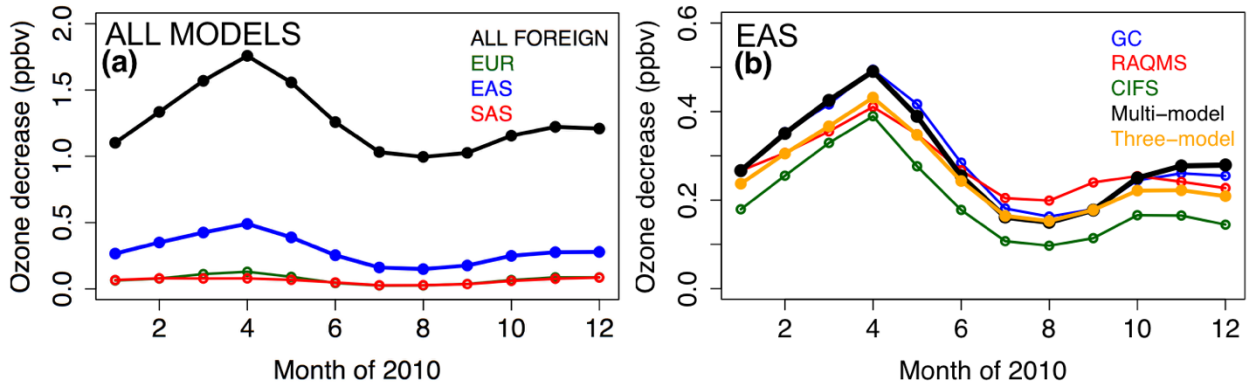


1490
 1491 **Figure 4.** Evaluation of the GEOS-Chem adjoint base NO₂ product (recorded at near the satellite
 1492 overpassing time) with the OMI NO₂ columns. The differences between OMI and GEOS-Chem
 1493 (OMI-modeled) tropospheric NO₂ columns ($\times 10^{15}$ molec./cm²) are shown for (a) May and (b) June
 1494 2010. Details of the comparison are included in Section 2.3.2.
 1495

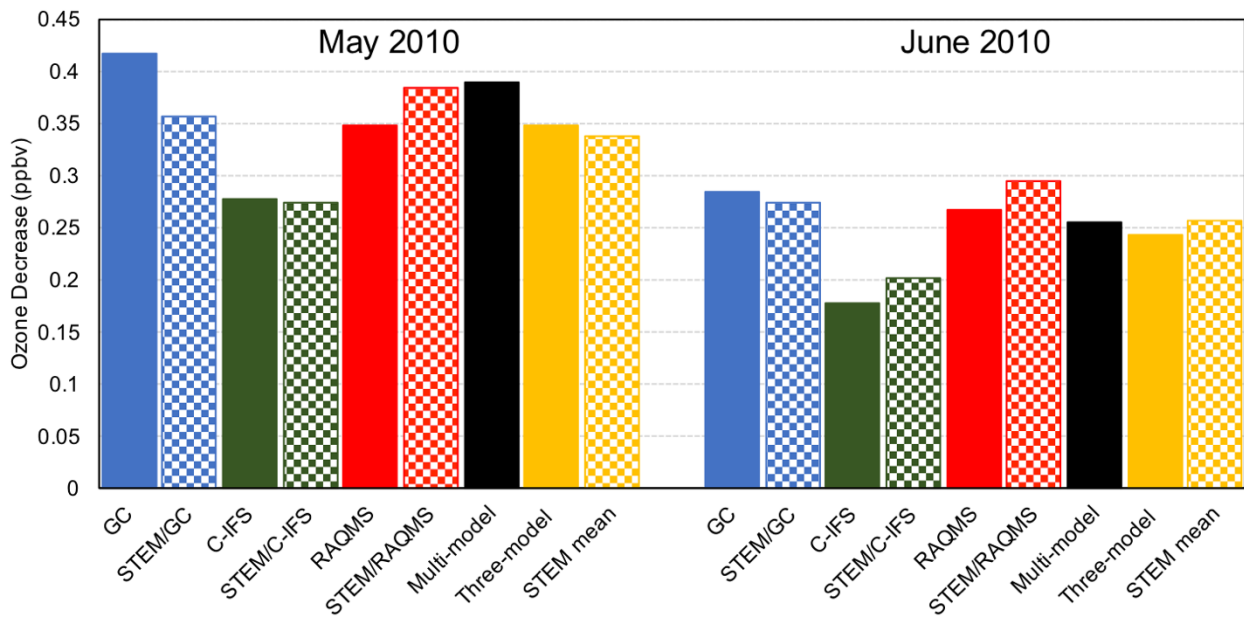


1496
1497
1498
1499
1500
1501
1502
1503

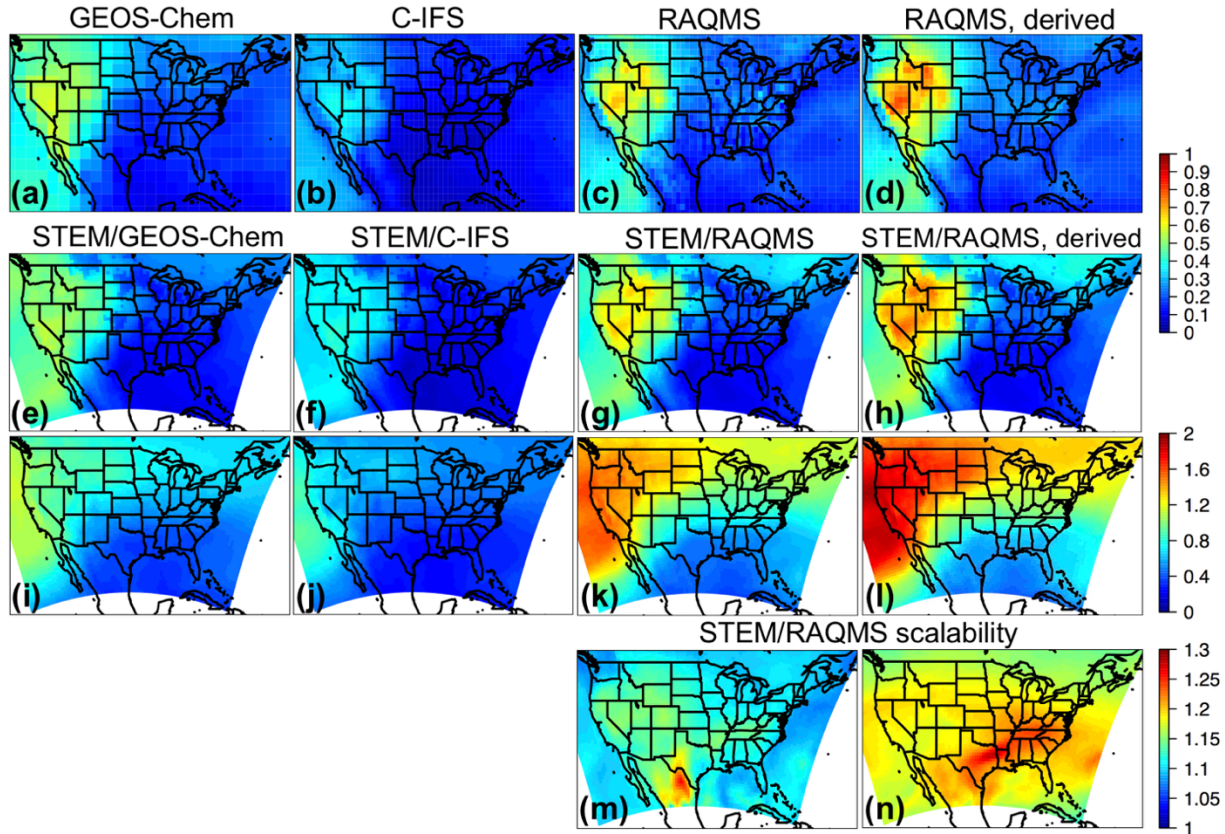
Figure 5. The RERER maps in May (left) and June (right) 2010 over the continental US, calculated based on the monthly mean O₃ from multiple global models' base and emission sensitivity simulations. The RERER metric (unitless) was defined in eq. (2) in the text. Values larger than 1 and smaller than 0 are shown in purple and grey, respectively. The US (including continental US as well as Hawaii which is not shown in the plots) mean values are indicated for each panel at the lower right corner. All models show declining RERER values from May to June, and the 7-model mean RERER values for May and June 2010 are ~0.5 and ~0.4, respectively.



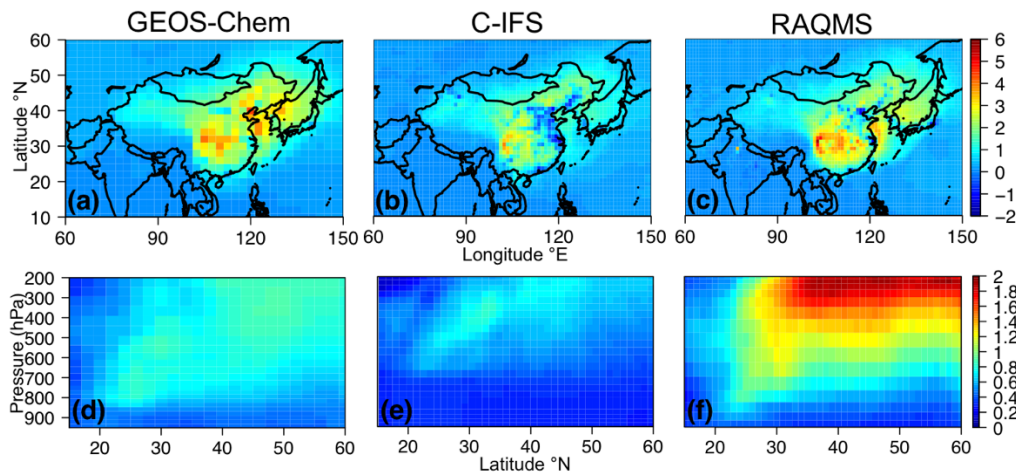
1504
 1505 **Figure 6. (a)** North American (130-65°W; 20-50°N) mean O₃ sensitivity to 20% anthropogenic
 1506 emission reductions in various non-North American regions, averaged from multiple (six-eight,
 1507 see details in text) global models. **(b)** North American surface R(O₃, EAS, 20%) values, as
 1508 estimated by single (the three STEM boundary condition models) or multi- global model means.
 1509 The “Multi-model” and “Three-model” in the legend indicate the mean sensitivities of all eight
 1510 global models and only of the three boundary condition models, respectively.
 1511



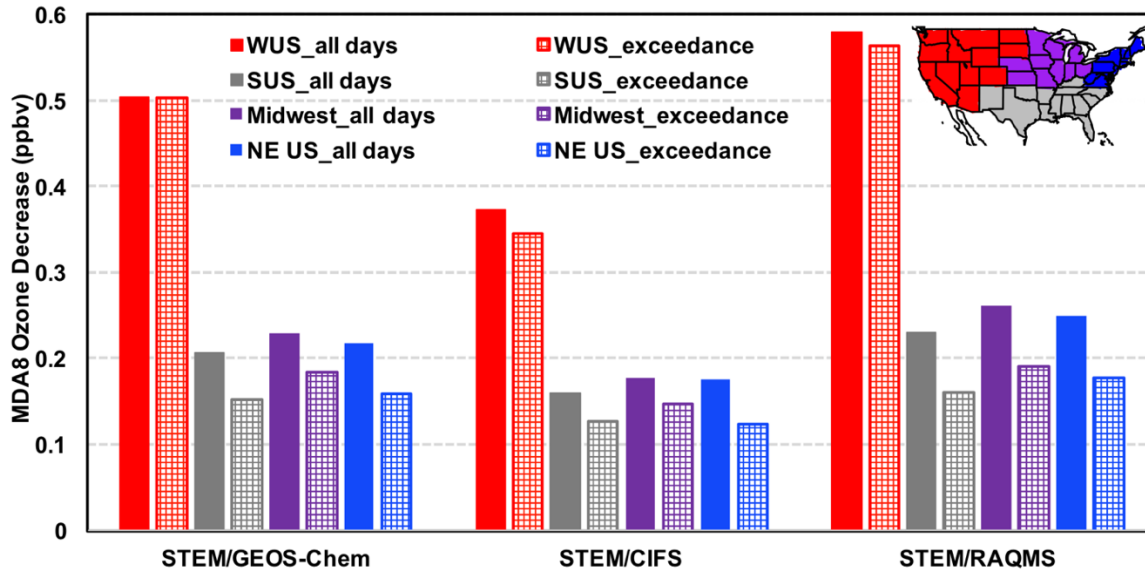
1512
 1513 **Figure 7.** Monthly-mean North American (130-65°W; 20-50°N) surface R(O₃, EAS, 20%) values
 1514 from multiple global and regional model simulations for May (left) and June (right) 2010. STEM
 1515 model mean values were calculated from its hourly output from 8 May and on. The “Multi-model”
 1516 and “Three-model” in the legend indicate the mean sensitivities of all eight global models and only
 1517 of the three boundary condition models, respectively.



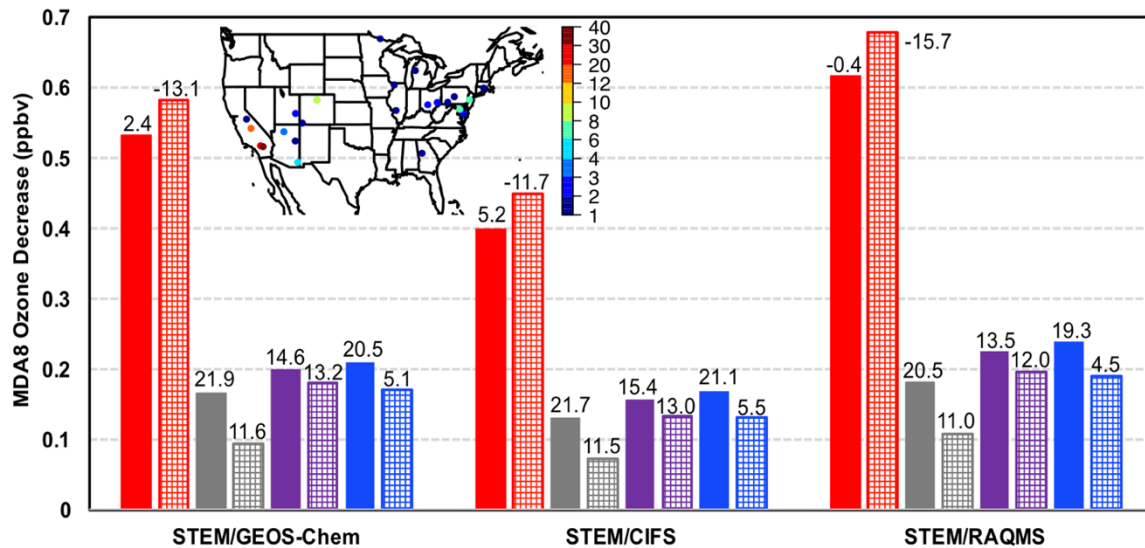
1518
 1519 **Figure 8.** The monthly-mean $R(\text{O}_3, \text{EAS}, 20\%)$ in June 2010 for: (a-d) surface O_3 (ppbv) from the
 1520 three boundary condition models, (e-h) STEM surface O_3 (ppbv), and (i-l) STEM column O_3
 1521 ($\times 10^{16}$ molecules/cm²). $R(\text{O}_3, \text{EAS}, 20\%)$ values from the simulations associated with GEOS-
 1522 Chem, ECMWF C-IFS, and RAQMS are shown in (a;e;i), (b;f;j) and (c;g;k), respectively. (d;h;l)
 1523 show 1/5 of the $R(\text{O}_3, \text{EAS}, 100\%)$ from the simulations related to RAQMS. STEM/RAQMS-
 1524 based “Scalability” S_{O_3} (eq. (3)) values over the NAM are shown for (m) surface and (n) column
 1525 O_3 .



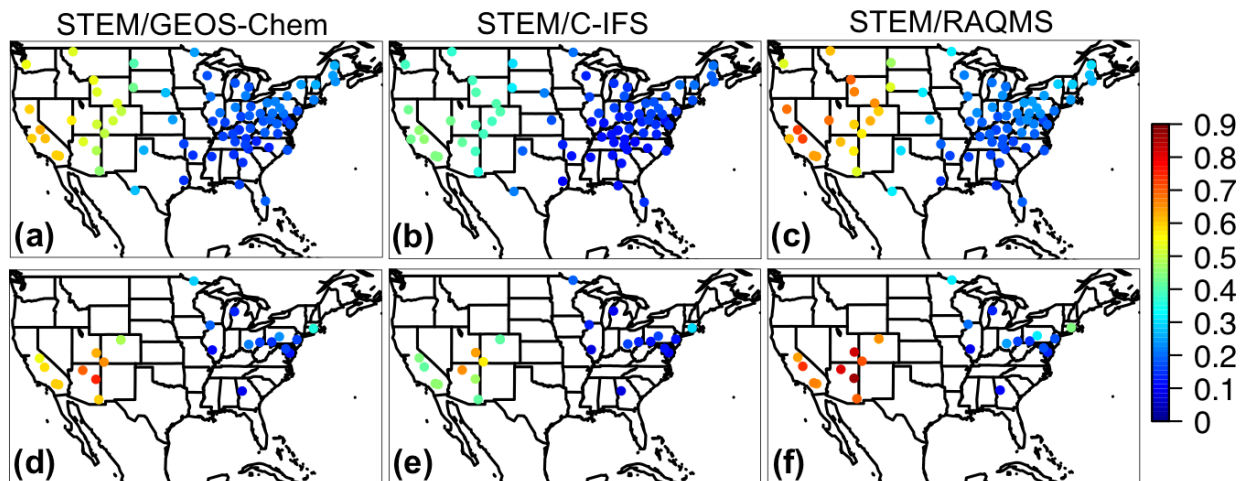
1526
 1527 **Figure 9.** The monthly-mean $R(\text{O}_3, \text{EAS}, 20\%)$ in ppbv in June 2010 from the three boundary
 1528 condition models at the source and near the receptor regions: (a-c) surface O_3 in the East Asia; and
 1529 (d) O_x (GEOS-Chem) or (e-f) O_3 (ECMWF C-IFS and RAQMS) along the cross section of 135°W
 1530 (near the west boundary of the STEM model domain as defined in Figure 2a).



1531
 1532 **Figure 10.** STEM R(MDA8, EAS, 20%) for May-June 2010 in four US subregions (defined in the
 1533 inset panel, also consistent with the definitions in Figures 2/S4 and Tables 2-3), averaged on all
 1534 days (bars with solid fill) and only on the days when the simulated total MDA8 O₃ concentrations
 1535 were over 70 ppbv (bars with grid pattern fill). The results from the STEM runs using GEOS-
 1536 Chem, ECMWF C-IFS and RAQMS boundary conditions are shown separately.
 1537

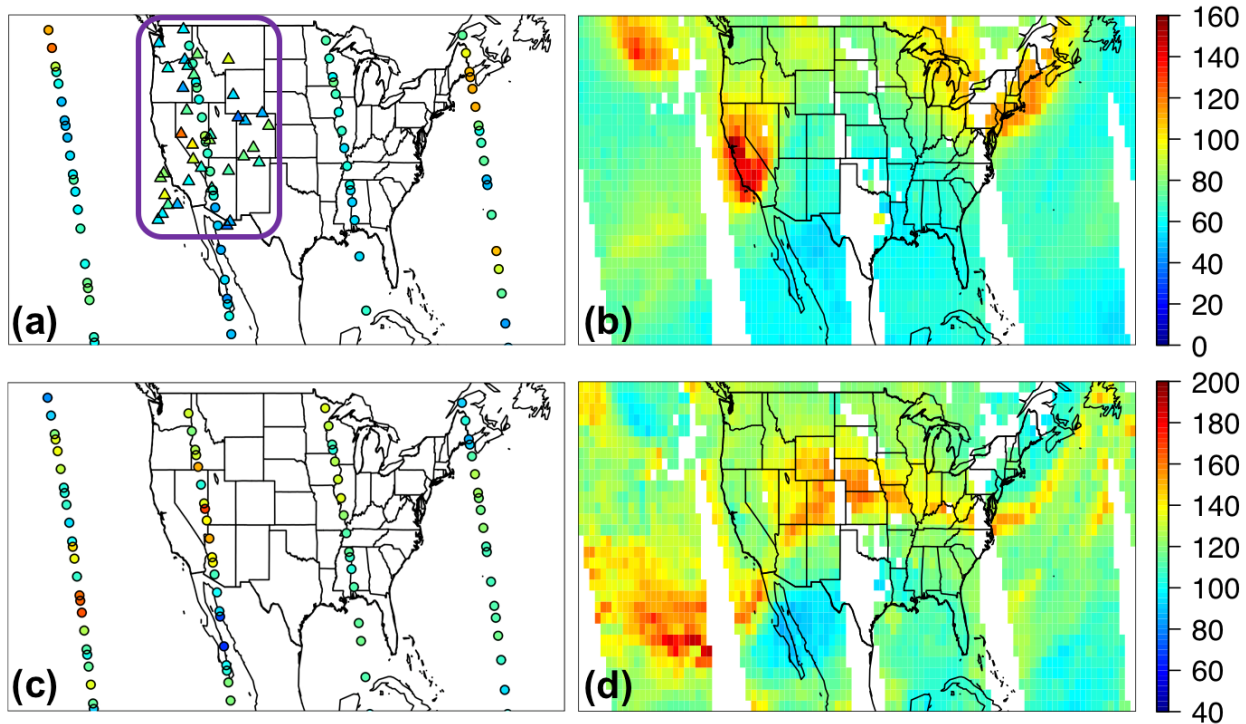


1538
 1539 **Figure 11.** STEM R(MDA8, EAS, 20%) for May-June 2010 at the CASTNET sites in four US
 1540 subregions (same definition as in Figure 10 inset), averaged on all days (bars with solid fill) and
 1541 only on the days when the observed MDA8 O₃ concentrations were over 70 ppbv (bars with grid
 1542 pattern fill). The results from the STEM runs using GEOS-Chem, ECMWF C-IFS and RAQMS
 1543 boundary conditions are shown separately. Biases for the corresponding model base runs are
 1544 shown above the bar plots. Inset shows at various CASTNET sites the number of days when the
 1545 observed MDA8 O₃ concentrations were over 70 ppbv.
 1546



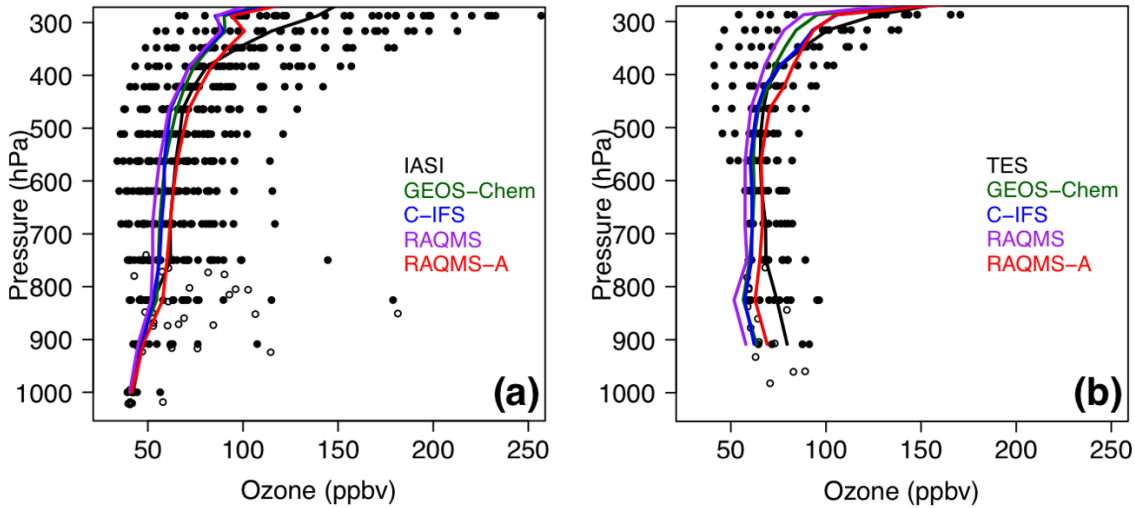
1547
1548
1549
1550
1551
1552

Figure 12. STEM R(MDA8, EAS, 20%) in ppbv for May-June 2010 at the CASTNET sites on (a-c) all days and (d-f) the days when the observed MDA8 O₃ concentrations were over 70 ppbv. The results from the STEM runs using (a;d) GEOS-Chem, (b;e) ECMWF C-IFS and (c;f) RAQMS boundary conditions are shown separately.

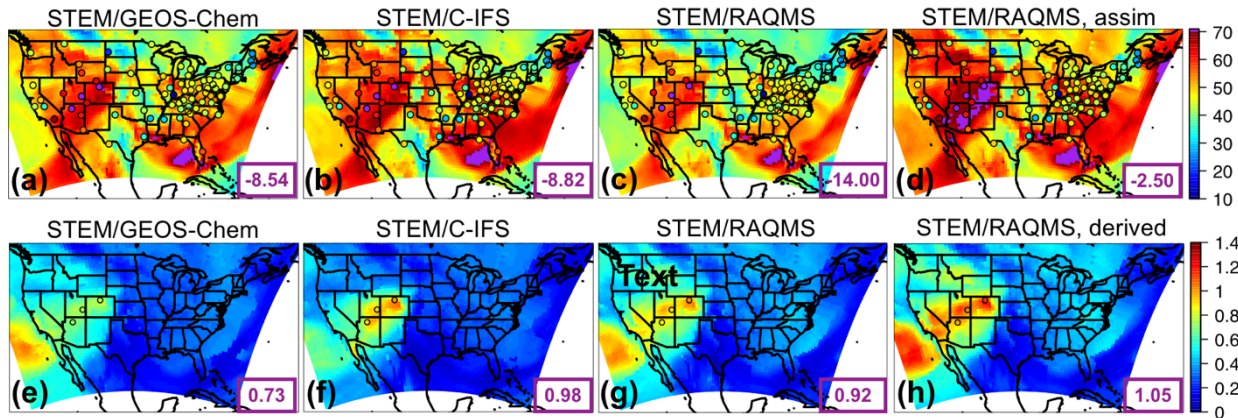


1553
1554
1555
1556
1557
1558

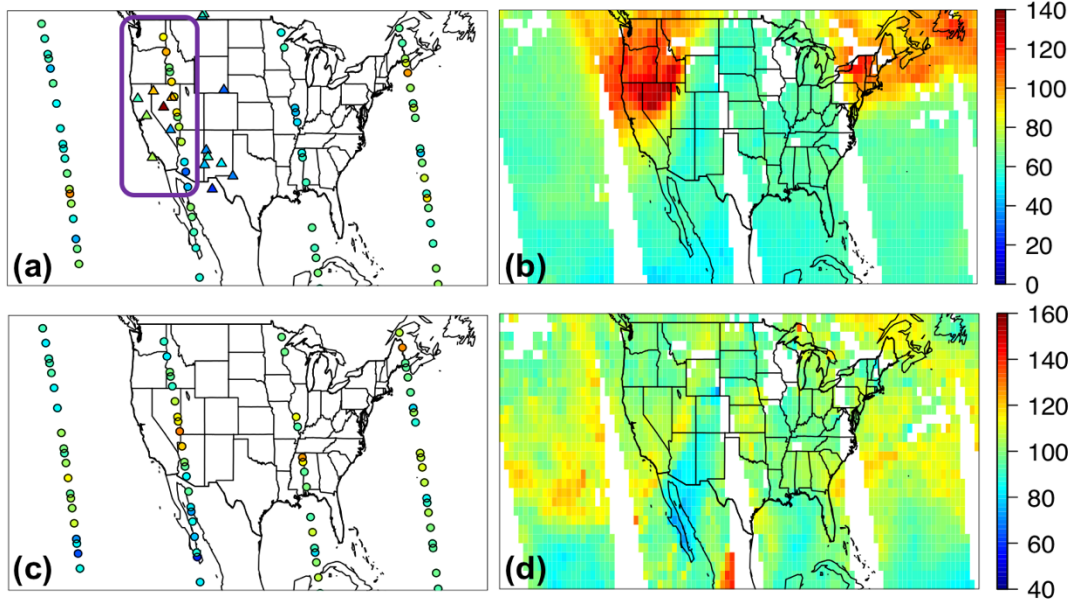
Figure 13. Case study of 9 May 2010: (a-b) Ozone (ppbv) and (c-d) CO (ppbv) at ~500 hPa from the L2 (a;c) TES retrievals (circles) and (b;d) L3 AIRS products at early afternoon local time. The L2 IASI O₃ (ppbv) at ~500 hPa retrieved using the TES algorithm (details in Section 2.3.2) at the mid-morning local times is shown on panel (b) as triangles. The O₃ profiles within the purple box in panel (a) were used in the model evaluation shown in Figure 14.



1559
 1560 **Figure 14.** Case study of 9 May 2010: The comparisons between (a) IASI and (b) TES O₃ in the
 1561 western US with the simulated O₃ in the STEM runs using the GEOS-Chem (green), C-IFS (blue),
 1562 RAQMS (purple), and assimilated RAQMS (red) boundary conditions. The O₃ profiles within the
 1563 purple box in Figure 10a were used in the evaluation. Observation operators were applied in the
 1564 comparisons (details in Section 2.3.2). Solid and open dots are TES/IASI data at the TES retrieval
 1565 reporting levels and at the variable surface pressure levels, respectively. Solid lines are median O₃
 1566 profiles from the satellite observations and the different STEM simulations, calculated only on the
 1567 TES retrieval reporting levels.
 1568

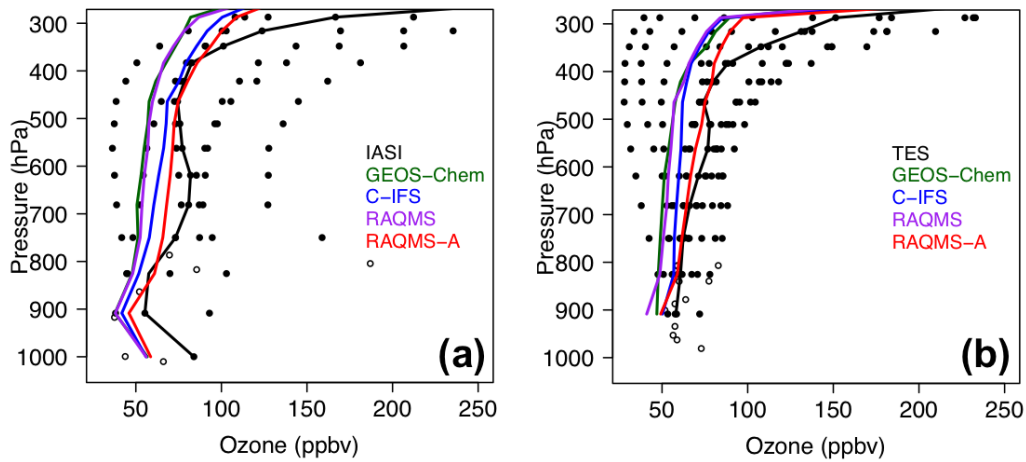


1569
 1570 **Figure 15.** Case study of 9 May 2010: (a-d) Surface MDA8 total O₃ and (e-h) surface R(MDA8,
 1571 EAS, 20%) from the STEM simulations using the (a;e) GEOS-Chem, (b;f) ECMWF C-IFS, and
 1572 (c;g) RAQMS free run as the boundary conditions. (d) Surface MDA8 total O₃ in a STEM base
 1573 simulation using the RAQMS assimilation run as the boundary conditions. CASTNET
 1574 observations are overlaid in filled circles in panels (a-d). (h) 1/5 of the surface R(MDA8, EAS,
 1575 100%) from STEM/RAQMS simulations. The conditions at ~400-500 hPa are shown in Figure S5.
 1576 Purple numbers at the lower right corners of (a-d) and (e-h) are mean model biases and mean
 1577 R(MDA8, EAS, 20%) values in ppbv at the three mountain sites (Grand Canyon NP, AZ;
 1578 Canyonlands NP, UT; and Rocky Mountain NP, CO) where O₃ exceedances were observed on this
 1579 day. The locations of these sites are shown in panel (e-h) as open circles.
 1580



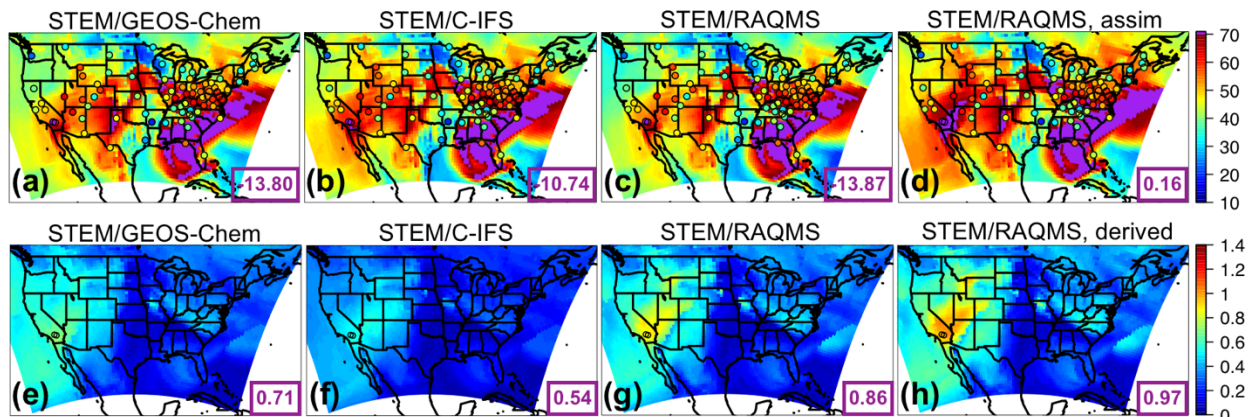
1581
1582
1583

Figure 16. Same as Figure 13, but for a case study of 10 June 2010.



1584
1585
1586

Figure 17. Same as Figure 14, but for a case study of 10 June 2010.



1587
1588
1589

Figure 18. Same as Figure 15, but for a case study of 10 June 2010. The CASTNET sites with O_3 exceedances on this day are Converse Station and Joshua Tree NP in southern California.

1590 **Table 1a.** HTAP2 base and sensitivity simulations by various global models. The STEM boundary
 1591 condition models are highlighted in bold.

Global model, Resolution: lon×lat×vertical layer, (References)	BASE	EASALL (-20%)	EASALL (-100%)	GLOALL (-20%)	NAMALL (-20%)	EURALL (-20%)	SASALL (-20%)
CAM-Chem, 2.5°×1.9°×56 (Tilmes et al., 2016)	✓	✓		✓	✓	✓	✓
CHASER T42, ~2.8°×2.8°×32 (Sudo et al., 2002)	✓	✓		✓	✓	✓	✓
EMEP rv48, 0.5°×0.5°×20 (Simpson et al., 2012)	✓	✓		✓	✓	✓	✓
SNU GEOS-Chem v9-01-03, 2°×2.5°×47 (Park et al., 2004; http://iek8wikis.iek.fz-juelich.de/HTAPWiki/WP2.3?action=AttachFile&do=view&target=_README_GEOS-Chem.pdf)	✓	✓		✓	✓		
CU-Boulder GEOS-Chem adjoint v35f, 2°×2.5°×47 (Henze et al., 2007)	✓	✓		✓	✓	✓	✓
RAQMS, 1°×1°×35, free running (Pierce et al., 2007, 2009)	✓	✓	✓				
RAQMS, 1°×1°×35, with satellite assimilation (Pierce et al., 2007, 2009)	✓						
OsloCTM3 v2, ~2.8°×2.8°×60 (Søvde et al., 2012)	✓	✓		✓	✓	✓	✓
ECMWF C-IFS, ~0.7°×0.7°×54/1.125°×1.125°×54, as the STEM chemical boundary conditions (Flemming et al., 2015)	✓	✓		✓	✓	✓	✓

1592 Acronyms:

1593 CAM-Chem: Community Atmosphere Model with Chemistry

1594 C-IFS: Composition-Integrated Forecasting System

1595 ECMWF: European Center for Medium range Weather Forecasting

1596 EMEP: European Monitoring and Evaluation Programme

1597 GEOS-Chem: Goddard Earth Observing System with Chemistry

1598 RAQMS: Realtime Air Quality Modeling System

1599 SNU: Seoul National University

1600 **Table 1b.** STEM regional simulations for HTAP2

Boundary condition model, Resolution: lon×lat×vertical layer	BASE	EASALL (-20%)	EASALL (-100%)
SNU GEOS-Chem v9-01-03, 2°×2.5°×47	✓	✓	
RAQMS, 1°×1°×35, free running	✓	✓	✓
RAQMS, 1°×1°×35, with satellite assimilation	✓		
ECMWF C-IFS, 1.125°×1.125°×54	✓	✓	

1601
1602 **Table 1c.** STEM and its boundary condition models' key inputs and chemical mechanisms, with
1603 references. More details on the models can be found in Table 1a and the text.

Model	Meteorology	Biogenic VOCs; NO _x	Lightning	Biomass Burning	Chemical Mechanism
GEOS- Chem	GEOS-5	MEGAN v2.1 (Guenther et al., 2012); Wang et al., 2009	based on GEOS-5 deep convective cloud top heights and climatological observations (Murray et al., 2012)	GFED v3.0 (van der Werf et al., 2010)	GEOS-Chem standard NO _x -O _x - hydrocarbon-aerosol (http://acmg.seas.harvard.edu/geos/doc/archive/man.v9-01-03/appendix_1.html)
RAQMS	Online (Pierce et al., 2007)				CB-IV (Gery et al., 1989) with adjustments
ECMWF C-IFS	IFS	MEGAN- MACC, (Sindelarova et al., 2014); POET database for 2000 (Granier et al., 2005)	based on IFS convective precipitation (Meijer et al., 2001)	GFAS v1.0 (Kaiser et al., 2012)	CB05 (Yarwood et al., 2005)
STEM	WRF-ARW v3.3.1	WRF- MEGAN v2.1	based on scaled WRF convective precipitation	FINN v1.0 (Wiedinmye r et al., 2011)	SAPRC99 (Carter, 2000)

1604 Acronyms:

1605 CB: Carbon Bond

1606 FINN: Fire INventory from NCAR

1607 GFAS: Global Fire Assimilation System

1608 GFED: Global Fire Emissions Database

1609 IFS: Integrated Forecasting System

1610 MACC: Monitoring Atmospheric Composition and Climate

1611 MEGAN: Model of Emissions of Gases and Aerosols from Nature

1612 POET: Precursors of Ozone and their Effects in the Troposphere

1613 WRF-ARW: Advanced Research Weather Research and Forecasting Model

1614 **Table 2a.** Evaluation of the period mean (1 May-30 June, 2010) multi- global model free
 1615 simulations against the CASTNET observations, only at the sites where 95% of the hourly O₃
 1616 observations are available. Evaluation of the individual models is summarized in Table 2b.

Subregion	US EPA regions contained	Number of sites	Mean bias (ppbv)		RMSE (ppbv)	
			3 BC ^a models	8 global models	3 BC models	8 global models
Western US	8, 9, 10	19	-5.68	-2.52	10.37	7.05
Southern US	4, 6	18	11.61	10.24	13.62	11.96
Midwest	5, 7	13	8.03	7.66	9.16	8.67
Northeast	1, 2, 3	17	9.55	10.63	10.28	11.24
All	1-10	67	5.49	6.22	11.11	9.96

1617 ^aBC: Boundary Conditions

1618
 1619 **Table 2b.** Evaluation of the period mean (May-June 2010) global model free simulations against
 1620 the EANET and CASTNET observations. The STEM boundary condition models are highlighted
 1621 in bold.

Network	Number of sites	RMSE (ppbv)							
		CAM-Chem	EMEP	CHASER	SNU GEOS- Chem	GEOS-Chem adjoint	RAQMS	OsloCTM3 v2	C-IFS
CASTNET	67	13.30	11.61	15.43	15.55	13.48	9.32	11.05	11.00
EANET	11	10.38	9.96	11.39	9.18	11.04	8.60	12.97	10.86

1622
 1623 **Table 2c.** Evaluation of the period mean (May-June 2010) multi- global model free simulations
 1624 against the EANET observations in Japan and Korea. Evaluation of the individual models is
 1625 summarized in Table 2b.

Country	Number of sites	Mean bias (ppbv)		RMSE (ppbv)	
		3 BC ^a models	8 global models	3 BC models	8 global models
Japan	8	0.36	1.01	8.77	9.25
Korea	3	1.14	3.98	8.37	10.51
All	11	0.57	1.82	8.66	9.61

1626 ^aBC: Boundary Conditions

1627 **Table 3a.** Evaluation of the hourly STEM simulated total O₃ (averaged from the three base
 1628 simulations that used the different free-running boundary conditions) against the CASTNET
 1629 surface observations for 8 May-30 June, 2010. The subregional mean R(O₃, EAS, 100%) and its
 1630 correlation coefficient with the observed O₃ are also shown.

Subregion	US EPA regions contained	Number of sites	Mean elevation (km): actual/model	Mean bias (ppbv)	RMSE (ppbv)	Correlation (model base; obs)	Correlation (obs; modeled EAS)	Mean EAS sensitivity (ppbv)
Western US	8, 9, 10	22	1.75/1.71	1.60	4.86	0.76	0.34	0.48
Southern US	4, 6	22	0.38/0.31	20.33	22.13	0.58	0.27	0.15
Midwest	5, 7	16	0.29/0.28	15.64	17.97	0.70	0.15	0.17
Northeast	1, 2, 3	20	0.36/0.26	20.94	24.16	0.47	0.17	0.21
All	1-10	80	0.73/0.68	16.17	18.30	0.66	0.13	0.20

1631 **Table 3b.** Evaluation of the hourly STEM simulated total O₃ (separately for three base simulations
 1632 that used the different free-running boundary conditions) against the CASTNET surface
 1633 observations for 8 May-30 June, 2010.
 1634

Subregion	US EPA regions contained	Number of sites	Mean bias (ppbv)/RMSE (ppbv)/Correlation (model base; obs)		
			SNU GEOS-Chem	C-IFS	RAQMS
Western US	8, 9, 10	22	1.68/4.83/0.77	4.16/6.63/0.70	-1.03/4.81/0.76
Southern US	4, 6	22	21.18/22.94/0.57	20.34/22.07/0.60	19.48/21.45/0.56
Midwest	5, 7	16	15.77/18.17/0.70	16.41/18.46/0.72	14.73/17.35/0.69
Northeast	1, 2, 3	20	21.25/24.36/0.47	21.86/24.80/0.48	19.71/23.40/0.45
All	1-10	80	16.57/18.62/0.66	16.89/18.84/0.67	15.03/17.52/0.64

1635

1636 **Table 4.** The ranges and standard deviations (ppbv, separated by “;”) of R(O₃, *source region*, 20%)
 1637 by 6-8 global models (defined in eq. (1a-d)), summarized by months in 2010. The monthly multi-
 1638 model mean values are shown in Figures 5-6.

Month/ Source region	All Foreign/ Non-NAM (ppbv)	EUR (ppbv)	EAS (ppbv)	SAS (ppbv)
Jan	0.38-1.69; 0.41	0.002-0.12; 0.05	0.02-0.72; 0.24	0.001-0.11; 0.04
Feb	0.92-2.07; 0.37	0.02-0.15; 0.05	0.16-0.91; 0.28	0.02-0.12; 0.04
Mar	1.30-2.37; 0.38	0.07-0.21; 0.06	0.24-1.03; 0.30	0.03-0.12; 0.03
Apr	1.42-2.46; 0.33	0.09-0.23; 0.05	0.33-1.07; 0.28	0.04-0.12; 0.03
May	1.24-1.91; 0.21	0.06-0.17; 0.04	0.24-0.75; 0.19	0.05-0.11; 0.02
Jun	1.03-1.41; 0.13	0.03-0.07; 0.02	0.14-0.39; 0.09	0.04-0.07; 0.01
Jul	0.86-1.18; 0.13	0.02-0.04; 0.01	0.08-0.22; 0.06	0.01-0.04; 0.01
Aug	0.80-1.19; 0.13	0.01-0.04; 0.01	0.07-0.20; 0.05	0.02-0.04; 0.01
Sep	0.85-1.18; 0.13	0.03-0.05; 0.01	0.10-0.25; 0.06	0.02-0.06; 0.01
Oct	0.96-1.31; 0.14	0.04-0.10; 0.02	0.17-0.42; 0.09	0.03-0.08; 0.02
Nov	0.90-1.48; 0.19	0.05-0.15; 0.04	0.17-0.54; 0.14	0.04-0.10; 0.02
Dec	0.73-1.67; 0.29	0.03-0.18; 0.05	0.14-0.66; 0.19	0.04-0.12; 0.03

1639