

Interactive comment on “Black carbon variability since preindustrial times in Eastern part of Europe reconstructed from Mt Elbrus, Caucasus ice cores” by Saehee Lim et al.

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We would like to thank the two anonymous referees for their careful reading of the manuscript, and also the time they dedicated to evaluating this study. All comments were highly insightful. Please find below our point-by-point response to the critiques and a highlight to the changes made to the manuscript to address these. We strongly feel that we were able to address all the points raised.

1. I know it is kind of tradition of the “Grenoble” group to classify the ice core data into summer and winter values and there might be circumstances where this is justified. However, this is always difficult, since the record does not contain clear time markers for the seasons and therefore assumptions have to be made. In this case, the 25th

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and 75th percentiles of thickness of an annual layer were arbitrarily chosen, assuming equal distribution of precipitation (and preservation on the glacier) throughout the year. This is conducted without explaining the hypothesis behind.

: We agree with the reviewer’s point that classifying the ice core data into summer (or warm season) and winter (or cold season) values was done without clear time markers. However, the ELB ice cores show very clear seasonality with the high summer values and the low winter values of water stable isotopic composition ($d_{18}O$ and dD) and NH_4^+ . Both $d_{18}O$ and NH_4^+ were thus used to classify seasonal ice layers of the cores (summer-half year and winter-half year). We think our proposed method for dating and seasonality is the best one for the moment. We also agree with the reviewer’s point that the 25th and 75th percentiles of thickness of a seasonal snow layer were chosen in this study assuming equal distribution of precipitation through a year and the hypothesis behind choosing this method is not clearly explained. In the Caucasus region, most of the annual precipitation occurs in the western and southern sections of the Caucasus, reaching 3240 mm y^{-1} at Achishkho weather station (1880 m). Precipitation ranges between 2000 and 2500 mm y^{-1} at 2500 m a.s.l. in the west and declines to 800–1150 mm y^{-1} in the east on the northern slope of the Caucasus (Mikhaleenko et al., 2015). A regular year round precipitation has been observed in the Western Caucasus at Klukhorskij Pereval station (2037 m a.s.l., 50 km westward; the location is indicated with number “7” in Kozachek et al., 2016, Fig. 1), where the proportions of mean summer and winter precipitations are 0.94 m (52%) and 0.87 m (48%), respectively, and precipitation of each month accounts for 6–11 % of total precipitation for the period 1966–2009 (www.meteo.ru). In our 2009 ELB core, the mean annual snow accumulation rate (1455 mm w.e. y^{-1} for the last 140 years) obtained by counting annual layers suggests that the deposited snow at the ELB site is well preserved without significant loss driven by wind erosion, although direct precipitation measurements are not available at the drilling site. Seasonal snow is also well preserved, with nearly equal deposition amounts from the warm season (45% of total accumulation) and the cold season (55% of total accumulation), e.g., a short firn core

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spanning the years 2012-2009 (Kutuzov et al., 2013). Hence, we suggest that it was reasonable to assume that precipitation at the ELB site is equally distributed through a year. Furthermore, our separation method of seasonal snow layers is supported by the coincidence of maximum (or minimum) values of both $\delta^{18}\text{O}$ and NH_4^+ in the annual snow layers. Most of maximum (or minimum) values of both $\delta^{18}\text{O}$ and NH_4^+ were observed in the 25th and 75th percentiles of thickness of a summer (winter) snow layer. However, we also observed unusual shift in NH_4^+ from $\delta^{18}\text{O}$ pattern, although it was rarely observed (roughly 10-year-ice layers over the entire 198-year-long record). The impact of inaccurate seasonal separation on rBC was limited by calculating median rBC mass seasonal concentration values. We finally concluded that the 25th and 75th percentiles of thickness of a seasonal snow layer were chosen in this study assuming equal distribution of precipitation through a year and that summer or winter rBC mass concentrations that were provided following this method are corresponding to summer maximum or winter minimum values of climate proxies in an annual layer.

We added a following sentence in line 150: "This seasonal separation method is fairly supported by the fact that (i) observed precipitation in the Western Caucasus is equally distributed throughout a year (e.g., at Pereval Klukhorskii observatory which is located at 2037 m a.s.l. and only 50 km from the drilling site: 52% of the annual precipitation (resp. 48%) is observed during summer (resp. winter) and each monthly precipitation accounts for 6-11 % of total precipitation for the period 1966-2009; www.meteo.ru) and (ii) maximum or minimum values of both $\delta^{18}\text{O}$ and NH_4^+ coincide for most of the Elbrus core annual ice layers."

2. The obtained summer and winter rBC records do show similar trends and the slight differences around 1900 are not discussed at all.

: As indicated by the reviewer, we observed slight increase of winter rBC values in 1900-1920 with respect to summer rBC values. We do not fully understand why the seasonal differences were shown in the period 1900-1920. There might be increased winter BC inputs for the period transporting to the ELB site. We used BC emission

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inventory data of Lamarque et al. (2010), which are well supported by rBC deposition reconstructed from Greenland ice cores (McConnell et al., 2007). In McConnell et al. (2007), rBC particles that transported from North America markedly increased at the beginning of 20th century, indicating increased BC emissions in North America for the period. The enhanced BC input from North America might be detected in the ELB ice core layers of the period 1900-1920, although it was not shown in our simulations.

The following sentence was added in line 254. "Meanwhile, the slight increase of winter rBC values in 1900-1920 with respect to summer rBC values are not well understood. Although speculative, it may reflect increased winter BC inputs transporting through the free troposphere (FT) from North America, where BC emissions markedly increased at the beginning of 20th century"

3. As expected the JJA and DJF scenarios for the atmospheric BC load are different, with much higher contributions from North America (NAM) to DJF. However, the authors do not question their summer and winter classification in the ice core because of this finding, but do instead explain the difference with an overestimation of the NAM footprint density by the simulation. To my opinion, the classification into JJA and DJF needs better justification, for example by showing that the annual ice core rBC concentrations and annual atmospheric BC loads agree less.

: Unlike scenario results for the seasonal atmospheric BC load at the ELB site, the seasonal rBC trends of the ELB ice cores were similar except for the periods 1900-1920 and 2000-2013. The winter rBC increased relatively for the first case and the summer rBC increased more obviously for the latter case. Both features were not shown in the simulations. Particularly, the increased winter rBC concentrations in ~1900-1920 may be linked with stronger BC inputs from North America at the beginning of the 20th century when the BC emissions were the strongest in that region. We do not understand well the mechanisms by which North American BC emissions could be strongly detected in the ELB ice core for this period. The following sentence was added in line 406. "Consequently, the observed overestimation of NAM contribution for winter

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at the ELB site (Fig 9b) is likely due to an overestimation of NAM footprint density in the statistical process applied on FLEXPART simulation data, although the stronger BC inputs from NAM might have contributed to the increased winter rBC concentrations of the ELB ice core at the beginning of 20th century.” Unfortunately, anthropogenic BC emissions from ACCMIP are available on the decadal scale only. We thus cannot show annual atmospheric BC load.

4. What is puzzling is that in the other manuscript about this ice core (Kozachek et al., CPD 2016), classification into seasons is conducted by introducing the mean $\delta^{18}O$ value as threshold. If at all, the procedure should be the same. Please also reconcile the details about the core (20.4 m here, 20.5 m in Kozachek et al.; dating uncertainty few years here, ± 1 year in Kozachek et al.) and include a reference to that manuscript.

: Mikhalenko et al. (2015) has established age scale of the ELB ice core using NH_4^+ and succinic acid, and posteriori validation with $d^{18}O$, resulting in a 2-year difference between annual layer counting of $d^{18}O$ signal and the NH_4^+ stratigraphy at 106.7 m. Kozachek et al. (2016) has made the ice core dating using annual $d^{18}O$, $d^{18}O$ threshold, and use of NH_4^+ and succinic acid if issues with $d^{18}O$. They initially reported a 1-year uncertainty of the dating but recently corrected this estimate to 2-years to agree with Mikhalenko et al. (2015) (see response to reviewers provided by Kozachek et al. (2016), TCD). Finally, both Mikhalenko et al. (2015) and Kozachek et al. (2016) have reported that a 2-year difference between annual layer counting of $d^{18}O$ signal and the NH_4^+ stratigraphy at 106.7 m. This is an excellent agreement on age scales that were obtained by two methods, suggesting robust dating results of the ELB ice core from top to 106.7 m. In our study, we discussed rBC annual variability down to 156.6 m, corresponding to year 1825. the dating uncertainty from the surface to 106.7 m is 2-years as indicated by Mikhalenko et al. (2015) and Kozachek et al. (2016), but the uncertainty may be larger below 106.7 m due to ice thinning.

We corrected in line 141: “The dating uncertainty is 2-year between 106.7 m and the top (Kozachek et al., 2016; Mikhalenko et al., 2015) and probably larger below 106.7

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m due to ice thinning” We corrected in line 72: “. a 20.5 m-long ice core (the 2013 core). . .”

5. The rBC size distribution data are very valuable since they support other findings that the rBC particle sizes in snow and ice are larger than in the atmosphere. However, the difference in MMD between summer and winter (Fig. 5) is not so obvious to me. The main discrepancy is for the few data points before 1960 where the data coverage is anyway poor. Have you tested if the MMDs for the period after 1960 are significantly different, considering the strong variability of the data?

: We agree with the reviewer’s point that the difference in MMD between summer and winter is not very obviously shown. We thus conducted student’s t-test on the MMD between summer and winter for the period 1960-2009 and the test resulted in significantly different mean MMD values between two seasons with $p < 0.01$. On the other hand, summer (or winter) MMDs for two periods for which highly variable rBC concentrations were observed, e.g., the period 1960-1999 and the period 2000-2009, were not significantly different. The following sentence was added in line 285: “No statistically significant temporal change in rBC MMD was identified over the 1940-2009 period.” The following sentence was added in line 302: “The difference in seasonal rBC size distributions are statistically significant ($p < 0.01$).”

6. Line 63: rephrase: that is reconstructed in the downstream of Europe.

: The sentence revised as follow: “The ice core record therefore provides information on long-term variability and evolution of BC emissions of Europe.”

7. Line 95: Please specify “upper section” (move this up from line 116).

: The “upper section” in the sentence was specified as follow: “The upper section of the 2009 firn core (surface to 7.2 m depth) was analyzed discretely.”

The firn depth in line 221 was revised to 7.2 m.

8. Line 117: Give max and min numbers of data points per year.

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: We added max and min numbers of data points per year in line 117: “The density of rBC data points per year ($N=8\sim 376$) depends on annual snow accumulation rates and ice thinning with depth.”

9. Lines 211-214: The fact that biomass burning emissions frequently occur in summer should be reflected in the emission estimates. I do not understand the argument for not considering the biomass emissions in DJF.

: We agree with the reviewer’s point that the argument why biomass burning emissions were considered only for summer simulations is not clear. ACCMIP inventory (Lamarque et al., 2010) provides anthropogenic BC emissions on a decadal scale and biomass burning (savanna and forest burnings) BC emissions on a monthly scale. The figure below (Fig. 2 here), which is a part of Saehee LIM’s PhD dissertation, shows that biomass burning BC emissions ($\text{kg}/\text{m}^2/\text{s}$) in Europe that were calculated by ACCMIP were frequent in summer time and minimized in winter time. The biomass burning BC emissions in May to August are larger by two orders of magnitude than those in November to February.

This is now clearly stated in the sentence in line 213 as follow: “We used anthropogenic emission only for constraining BC emissions in DJF and both anthropogenic and biomass burning emissions for constraining BC emissions in JJA, because seasonal biomass burning BC emissions are maximized in summer time (May to August), being two orders of magnitude larger than during winter time (September to February), as respect to anthropogenic emissions occurring year-round (Lamarque et al., 2010).”

10. Lines 232-236: Include Jungfraujoch (e.g. Bukowiecki et al., Aerosol and Air Quality Research, 16: 764–788, 2016).

: The reference (Bukowiecki et al., 2016) was added with relevant discussion in lines 232-236. In line 231, the following sentence is added. “In contrast to the boundary layer sites, the atmospheric measurements at high-elevation sites in Europe (e.g., Puy de Dôme at 1465 m a.s.l., Sonnblick at 3106 m a.s.l. and Jungfraujoch at 3580 m a.s.l.)

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revealed 2 to 3 times higher EC levels during summer than winter (Bukowiecki et al., 2016; Pio et al., 2007; Venzac et al., 2009), ”

11. Line 293: Mikhalenko et al. (2015) do not mention aerosol removal processes. Please clarify that you assume that wet deposition dominates, since there is often and regular precipitation throughout the year.

: We agree with the reviewer’s point that Mikhalenko et al. (2015) did not mention aerosol removal processes and the line 293 in our manuscript can be misleading. We therefore corrected the sentence as follow: “The shift of rBC sizes induced by dry deposition should be negligible, as quite high (100-200 mm/month) and fairly constant precipitation rate throughout the year near the drilling site (e.g., 52% and 48% of annual precipitation observed in summer and winter, respectively, at Klukhorskii Pereval station (2037 m a.s.l., 50 km westward; Kozachek et al., 2016) suggests that wet deposition can be the dominant aerosol removal pathway at this site.”

12. Line 355: Matthias, 2004: Is this Matthias and Bösenberg, 2002?

: The reference “Matthias, 2004” in line 355 should be “(Matthias et al., 2004)” and the relevant reference info should be revised. Matthias et al. (2004) showed regular lidar observations of the vertical aerosol distribution at 10 European Aerosol Research Lidar Network (EARLINET) stations since 2000, for which they used the planetary boundary layer (PBL) height (km asl) at each station. Two mountain stations (L’Aquila at 1742 m agl and Potenza at 1536 m agl) showed monthly mean PBL height above 2 km asl and often higher weakly PBL height up to 3 km asl, while the PBL height was lower with monthly mean PBL of 1-2 km asl at the other stations. Thus, simulations for summer particle footprint within the lower 2 km layer in the atmosphere performed in our study are fairly consistent to the real PBL height at an area surrounding mountain and realistic aerosol transport to the drilling site.

13. Figure 2 would benefit from a better quality map. Please indicate location of ELB and explain abbreviations in the figure (NAM etc).

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: Figure 2 (Fig 3. here) was replaced with a better quality map as follow. The location of ELB is indicated by a red circle symbol. The abbreviations were explained in the final version of the manuscript as below, but not in this comment due to limited space for figure caption.

Figure 2. Five sub-regions classified as potential rBC emission source regions. Elbrus drilling site (43°20'53,9"N, 42°25'36,0"E) is indicated by a red circle. WEU, CEU, EEU, NAF and NAM represent Western Europe, Central Europe, Eastern Europe, North Africa and North America, respectively.

14. Fig. S1: The overlap between the 2009 and 2013 cores is not convincing. Could you support this with other ice core parameters (e.g. stable isotopes)?

:An overlapping section (m w.e.) of the 2013 core and the 2009 core was described with water stable isotope (d18O) values in the following figure. The common d18O values were observed in the 2013-core depth of 6.8-10.7 m w.e., corresponding to year 2009-2007. The current Fig. S1 was replaced with the following figure (Fig 4 here).

Figure S1. An overlapping section of the 2009 core and the 2013 core. We used the common d18O feature dated year 2009-2007 and located at 7-11 m w.e. depth along the 2013-core depth scale to extend the 2009-core record (main core) with the 2013-core record.

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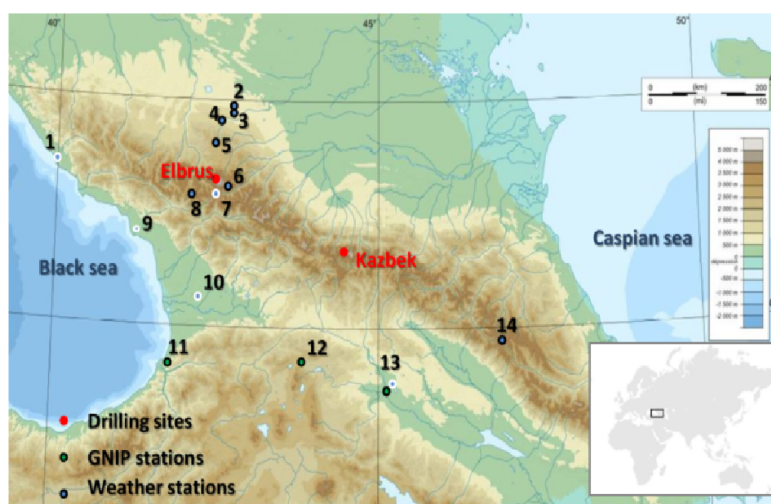


Fig. 1. Map showing the region around Elbrus (black rectangle in the world's map in the lower right corner). Klukhorskij Pereval station is indicated with number "7". (from Kozachek et al., 2016).

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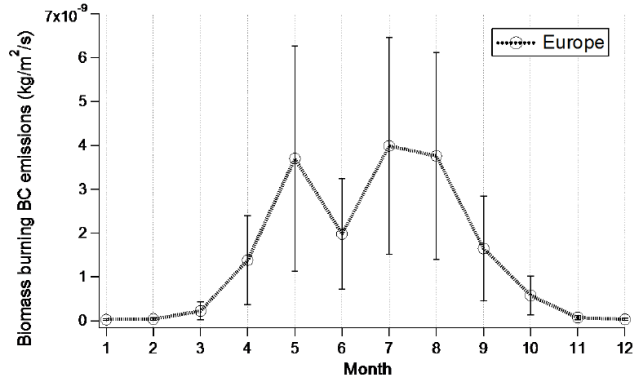


Figure I-10. Seasonality of BC emissions from biomass burning in Europe. The biomass burning BC emissions here include emissions from savanna burning and forest fires during the period 1900 to 2000. Mean monthly BC emissions with one standard deviation are estimated from ACCMIP BC emission inventory. Sources are from Lamarque et al. (2010).

Fig. 2. Seasonality of BC emissions from biomass burning in Europe during the period 1900 to 2000. Sources from Lamarque et al. (2010).

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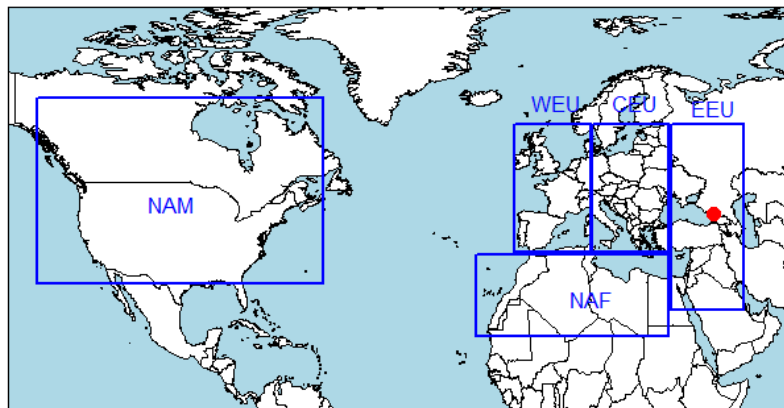


Fig. 3. Five sub-regions classified as potential rBC emission source regions. Elbrus drilling site (43°20'53,9"N, 42°25'36,0"E) is indicated by a red circle.

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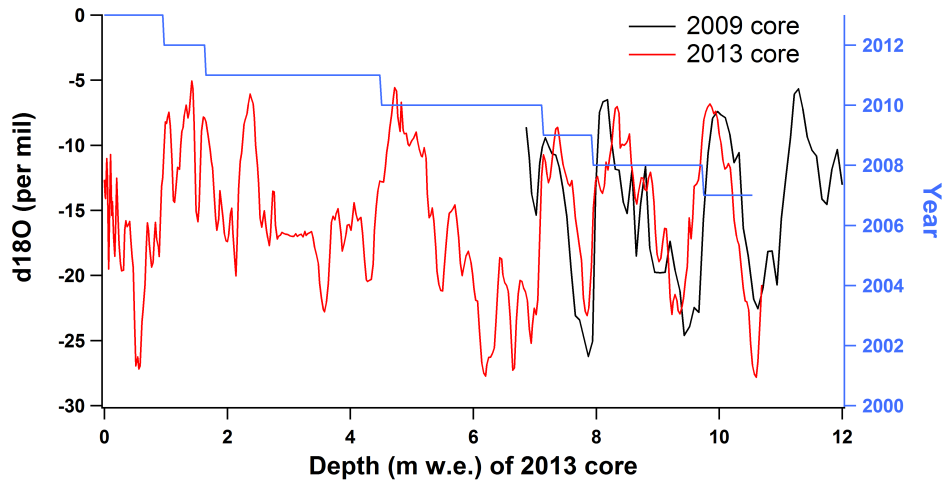


Fig. 4. An overlapping section of the 2009 core and the 2013 core.