

Author response to the ACPD manuscript: “ α -Pinene secondary organic aerosol yields increase at higher relative humidity and low NO_x conditions” by Lisa Stirnweis et al.

L. Stirnweis et al.

The authors would like to thank the two anonymous reviewers for their helpful and thoughtful comments and suggestions. Based on their comments we have modified the text, especially the parts related to the model description, NO_x regime and vapour wall losses, where additional explanations seemed to be necessary. Each specific point of the comments is addressed below. Reviewer comments are in italic typeset, our responses are in regular typeset and changes in the manuscript are highlighted in blue. All references are listed at the end of the author response. The author response is followed by a marked-up manuscript version showing the changes made.

Anonymous Referee #1

This work is laboratory and modeling study of SOA from alpha-pinene photooxidation under different RH conditions, and various NO_x/VOC ratios. The authors investigated the effects of various seed compositions, notably hydrophobic vs. hydrophilic seed, which in my opinion is a clever way to deduce that the RH effect is significantly based on liquid water and particle phase mixtures. The authors primarily used AMS, PTRMS, and SMPS to measure the compounds/particles of interest and performed wall loss corrections for particles under one RH condition. Vapor wall loss was not considered. Appropriate blanks were performed. The work demonstrates a liquid-water-based enhancement of SOA yields that may be due to a combination of many chemical and physical factors. The authors also attempted to give insight into phase partition by using the AOIMFAC model and their own observational inputs. The paper is well-written and the method is thoroughly described. A more thorough discussion of the mechanisms involved, and additional clarity about the modeling, would be welcomed. I have some comments and suggestions before publication can be recommended.

General comments:

1. The authors stated that lack of corrections for vapor wall deposition do not “influence the comparison between the experiments.” The statement is hard to understand when the RH dependence of vapor wall loss has been documented. For example, please see Loza et al EST (2010) and Nguyen et al PCCP (2016), where hydroperoxide, hydroxyepoxide, and organic

acid wall losses were measured under different humidity conditions in chambers and differ substantially between different RH conditions. Nguyen et al PCCP (2016) even gave a parameterization for these compounds as a function of RH in a 24 cubic meter Teflon chamber (e.g., $k_{wall_HMHP} = -1.4 \times 10^{-5} \times RH \text{ min}^{-1}$, $k_{wall_H2O2} = -9.6 \times 10^{-6} \times RH \text{ min}^{-1}$, and $k_{wall_HCOOH} = -2.2 \times 10^{-6} \times RH \text{ min}^{-1}$). As the authors can see, not only is vapor wall loss different for each RH condition, it is different for each chemical compound. The papers listed to support the authors' statement, namely Zhang et al PNAS (2014) and Nah et al ACP (2016), were only studied under dry ($RH < 5\%$) conditions so are not applicable to the current case. I do not believe that retro-actively applying the vapor wall loss corrections is critical to this work, but request that the authors conservatively estimate the errors that ignoring such a correction in the alpha-pinene system (which is known to produce compounds readily lost to walls) would cause. This may actually increase the enhancement that the authors observed.

Response: We have already acknowledged in the original version of the manuscript that vapour wall losses may influence chamber results and complicate data interpretation. This is especially the case if losses depend on the RH, as these losses would lead to biases that were not yet, but will be discussed in the manuscript. Nevertheless, we note that losses of the small organic molecules reported in Loza et al. (2010) and Nguyen et al. (2016), e.g. glyoxal, epoxides and peroxides are very likely due to their reactive uptake, the rate of which changes with RH. These processes occur at time scales of hours, much longer than those related to the absorption of SOA forming semi-volatile compounds onto the clean Teflon walls. The dependence of vapour absorption on RH (due to a change in wall accommodation coefficients or in the activity of the wall absorbed compounds) has not been reported to the best of our knowledge and indeed merits further investigations that are beyond the scope of the current study. Nevertheless, we believe that vapour absorption onto the Teflon walls is unlikely to be affected by RH, due to the hydrophobic nature of Teflon and its minor interactions with water at relatively low RH (<75%).

Based on the reviewer suggestions, we have discussed these effects more thoroughly in the revised version of the manuscript:

Eq. (8) does not take into consideration the loss of SOA-forming vapours onto the clean Teflon walls, which may suppress SOA yields from laboratory chambers under certain conditions. These processes may be related to the vapours' reactive uptake onto walls (Loza et al., 2010; Nguyen et al., 2016) or to their absorptive uptake (Zhang et al., 2014; Nah et al.,

2016). The reactive uptake of organic vapours onto chamber walls is only significant at high RH. Loss rates for important reactive gases, including glyoxal, epoxides and peroxides have been documented at different RH (Loza et al., 2010; Nguyen et al., 2016). While these processes may also influence reactive SOA-forming compounds under our conditions, they occur at time-scales of hours (Nguyen et al., 2016), much longer compared to the time scales of the absorptive uptake, e.g. based on recent direct measurements of vapour losses onto Teflon walls (~10 min (Krechmer et al., 2016)).

The absorption of organic compounds onto the chamber walls obeys Henry's law and depends on the compounds' accommodation coefficients and their activity at the wall/gas interface (see for example Zhang et al., 2015). The dependence of the compounds' absorption on RH (due to a change in accommodation coefficients or in the activity of the wall absorbed compounds) has not yet been reported to the best of our knowledge and indeed merits further investigations that are beyond the scope of the current study. Nevertheless, we believe that vapour absorption onto the walls is unlikely to be significantly affected by RH, due to the hydrophobic nature of Teflon and its minor interaction with water under subsaturation conditions (RH < 80 %). In our case, we have maintained chamber conditions during our experiments such that vapour wall losses and their inter-experimental differences can be minimized as much as possible. This is done by (1) maintaining a relatively constant wall-to-seed surface ratio for all experiments to avoid systematic biases between experiments and (2) increasing SOA production rates, which rapidly provide a significant particle condensational sink into which condensable gases can partition. We also note that vapour wall losses were found to be minor for the α -pinene SOA system where SOA formation is dominated by quasi-equilibrium growth (Zhang et al., 2014; Nah et al., 2016).

2. Adding to that subject, the authors are also suggested to monitor particulate wall losses at different RH conditions and for different composition in their future works in lieu of picking an average and going with it for all particles and all conditions. There is usually a noticeable difference in the rates of deposition depending on particle characteristics and wall wetness (related to RH). This is especially advisable since the authors are using so many different seeds that may respond to wall wetness in different ways (I would guess the CF seeds would not show the same increase in sticking as the hydrophilic seeds when water layers on the walls increase, with similar effect to their SOA-derived observations). Again, the suggestion here is to approximate and report uncertainty that may be caused by ignoring these dependencies in the revised version of the paper.

Response: Unlike vapours, particles are lost to the walls irreversibly and with an accommodation coefficient of unity. Therefore, particle wall losses, driven by diffusion and gravitation, are less dependent on the wall wetness (only through a change in the Teflon electrostatic properties with RH), and are strongly a function of particle diameters. As the organic particles and inorganic seeds are externally mixed and have different diameters, we have chosen to derive the loss rates of the organic phase by fitting its decay with time (instead of using the seed decay), when OA production is expected to be negligible. We have derived a loss rate for each experiment, which was used for wall loss correction. An average rate was only used in the case of insufficient statistics to perform accurate fitting. Therefore, we believe that our wall loss procedure takes into account the experiment-to-experiment variability.

3. I have a general criticism of the way “NO_x” is used in this paper. It is too vague. When the authors say “44.4(0.8) ppbv of NO_x” does that mean 42 ppb of NO₂ and 2.4 ppb of NO or any of the other innumerable combinations. . .? Additionally, it’s not really “NO_x” that’s important here, but rather nitric oxide (NO) because it changes the course of the reaction with the RO₂ radical, while NO₂ doesn’t do very much unless the precursor is an aldehyde (which in this case it is not). Can the authors be more clear about how much NO there is, instead of “NO_x”?

Response: Based on our previous calculations (see Platt et al. (2014)), RO₂ radicals would predominantly react with NO, when the concentration of the latter is higher than only 1 ppb. These conditions can be considered as high NO_x.

As mentioned in the manuscript, during high NO_x experiments, before lights were turned on, we have injected equal amounts of NO and NO₂, resulting in initial NO_x concentrations between 20-75 ppb. During these experiments and throughout the period when the majority of α-pinene was consumed, the NO concentration remained higher than 5ppb, indicating that α-pinene oxidation proceeded under high NO_x.

During low NO_x experiments, the average NO_x concentration was around 1-2 ppb, predominated by NO₂ while the NO concentration was below detection limits (<0.1ppb). Under these conditions RO₂-RO₂ reactions may prevail.

Based on the reviewer comment, we have updated the text on page 5, as follows:

Within the results section, the two terms “low NO_x” and “high NO_x” refer to the following conditions:

- Low $\text{NO}_x = \text{NO}_x/\alpha\text{-pinene} < 0.1$, with continuous HONO injection, indicated by an asterisk in Table 1 and figures.

- High $\text{NO}_x = \text{NO}_x/\alpha\text{-pinene} > 1$: Initial injection of NO + NO₂ with continuous HONO injection.

In Table 1, we report for high NO_x conditions the initial NO_x concentration, (which decays with time) and for low NO_x conditions the mean NO_x concentration. We note that the NO levels are the main driver for determining whether RO₂-NO or RO₂-RO₂ reactions would prevail. Based on our calculations (not shown here, see Platt et al., 2014), RO₂ radicals would predominantly react with NO, when the concentration of the latter is higher than only 1 ppb. These conditions can be considered as high NO_x . During high NO_x experiments and throughout the period when the majority of α -pinene was consumed, the NO concentration remained higher than 5ppb, indicating that α -pinene oxidation proceeded under high NO_x . During low NO_x experiments, the average NO_x concentration was around 1-2 ppb, predominated by NO₂, while the NO concentrations were below detection limits (<0.1ppb). Under these conditions RO₂-RO₂ reactions may prevail.

4. The authors mentioned hydrophilicity and solubility several times in the article, yet it's not clear how this is considered by the model, if at all? Also, despite the authors' statement that the few products considered in the model are adequate, more support is needed to understand how these few products can be fully representative of such a complex chemical system.

Response: The aim of the modelling part is to investigate whether thermodynamics alone can account for the higher yields observed at higher RH. The compounds used for the simulations were chosen from previously identified photo-oxidation products of α -pinene. The partitioning of these compounds between the gas and the condensed phases depends on their effective saturation concentrations, which is defined in Eq. 11. This equation takes solubility and hydrophilicity into account because it does not depend only on the vapour pressures of the pure compounds but also on their activity coefficients in solution (i.e. affinity towards other compounds in the organic phase, affinity towards water and affinity towards the inorganic species). We use AIOMFAC to calculate activity coefficients and perform a phase equilibration calculation between the gas and condensed phases. With this approach, we fully account for solubility (in the organic phase) and hydrophilicity (affinity towards particulate water).

As we have mentioned in the manuscript, the choice of the model compounds is constrained by the experimentally determined volatility distributions and O:C ratios. This significantly

constrains the nature of the compounds in a given volatility bin. For example, compounds with saturation concentrations of 100-1000 $\mu\text{g m}^{-3}$ would need to have a short carbon backbone chain (carbon number ~ 5), if their O:C ratios should match the observed O:C ratios of ~ 0.6 . The activity coefficients calculated by the model are more sensitive to the chosen compounds O:C ratios than to the compounds' structure and functional groups.

From the comments/questions of both reviewers, further details and clarifications seem to be necessary for a better explanation of the procedure adopted for the phase partitioning calculations. Therefore, in the corrected version of the manuscript, we have significantly modified section 2.6, relative to phase partitioning calculations.

Page 10, lines 12-29 were replaced by the following:

2.6 Thermodynamic modelling

General principles. Using thermodynamic modelling, we seek to understand whether observed changes in SOA bulk properties (yields and O:C ratios) with RH can be explained by a change in the particles' thermodynamic properties. Gas-liquid and liquid-liquid phase partitioning calculations are performed following the methods developed in Zuend et al. (2008; 2010; 2012) and Zuend and Seinfeld (2013), using the thermodynamic group-contribution model, AIOMFAC (Aerosol Inorganic-Organic Mixtures Functional groups Activity Coefficients) to calculate activity coefficients. This approach enables predicting the phase partitioning of known organic compounds knowing their abundances in a given known mixture of organic species and electrolytes at a given RH and temperature.

The modelling requires the use of explicit surrogate compounds. Based on Eq. (10), the partitioning of these compounds is driven by their volatility distributions, which depends on (1) the compounds' effective saturation concentrations and (2) their relative abundances. Compounds' effective saturation concentrations, used as model inputs, were calculated based on Eq. (11) initially assuming ideal mixing ($\gamma=1$) and utilizing vapor pressures estimated using EVAPORATION (Compernelle et al., 2011). The relative abundances of these compounds in the model mixtures are based on the volatility distributions derived from experimental data (section 2.5).

We assumed instantaneous reversible absorptive equilibrium of semi-volatile organic species into ideal and non-ideal liquid phase aerosols. In the case of non-ideal solutions, positive and negative deviations of mole fraction-based activity coefficients from unity indicate the degree of non-ideality in a mixture. The activity coefficients take into account the compounds'

affinity towards the solution (interactions with other organic species, electrolytes and water) and hence depend on the solution chemical composition. Therefore, the activity coefficients for the different organic species cannot be set *a priori*, but are calculated iteratively in the model (until convergence of the compounds' abundances in the different phases; see (Zuend et al., 2008; Zuend et al., 2011)).

We considered cases with and without interactions between the electrolyte and the organic phases, which enables assessing the solubility of the organic compounds in the inorganic seed aerosols. Seed concentrations in $\mu\text{g m}^{-3}$ were transformed in moles of seed per volume (mol m^{-3}) assuming equal shares of AHS and SA for acidic seeds and equal shares of AHS and AS for neutral seeds. Because interaction parameters of some organic functional groups with HSO_4^- are missing in AIOMFAC, we assumed the interactions of organic compounds with HSO_4^- and with SO_4^{2-} to be similar. For all computations, metastable supersaturated salt solutions were allowed.

Model compounds are only surrogates. The choice of these surrogates should reflect the wide range of volatility and hydrophilicity of SOA species. We have selected as surrogates α -pinene photo-oxidation products reported in literature, identified under different NO_x and aerosol seed conditions, covering the wide range of volatility relevant to SOA. Compound hydrophilicity instead depends heavily on the number of functional groups present in the molecule, which can be largely simulated by the compound O:C ratio (Zuend and Seinfeld, 2012). As increasing the compound O:C ratio also decreases its volatility, we have considered cases with and without fragmentation products with shorter carbon backbone chain, but high O:C ratios, expected to be representative of later generation photolysis and photo-oxidation products (Mutzel et al., 2015; Krapf et al., 2016). This approach would effectively decouple compounds' hydrophilicity and volatility and allow scanning both properties independently. While we recognize that oxidation conditions, e.g. NO_x concentrations, may significantly alter the product distribution, we did not select different sets of products for the different conditions. This is because:

- (1) such separation would implicitly suggest that the chemical composition of the few compounds reported at different conditions can be extrapolated to the bulk OA under our conditions;
- (2) such separation would significantly limit the number of surrogates at each condition, increasing the sensitivity of the model to the compounds' selection;

(3) the model is less sensitive to compounds' chemical structure than to their elemental composition (see below and Li et al. (2016)), e.g. number of oxygen and carbon, which has been taken into account by including fragmentation products.

The relative abundance of the selected compounds is optimized in the model for each experiment at the prevailing RH such that the modelled and measured SOA yield and O:C ratio match. Then, the RH is modified in the model, and changes in the SOA yields and O:C ratios are compared to the observations. Model calculations were performed at an OH exposure of $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$. In the following we thoroughly describe the different steps involved in the model setting.

Some detailed comments

Pg 13, ln 12: It's not clear why the authors conclude that the NO_x dependence is definitely due to low NO_x conditions forming less volatile compounds? Perhaps the low- NO_x conditions form more soluble compounds? Perhaps more well-mixed particles?

Response: Conclusions on the NO_x dependence are based on the multilinear analysis results in section 3.1. We have observed that independent of the prevailing RH, yields obtained under low NO_x are significantly higher than those obtained under high NO_x conditions. This implies that independent of the compounds' affinity towards water, compounds formed under low NO_x are less volatile than those obtained under high NO_x and tend to remain in the particle phase.

Pg 13, ln 18: How much of the reaction is actually ozonolysis? The authors should give an indication of ozone mixing ratio in these reactions, and calculate the prevalence of side reaction given the O₃+ α -pinene rates. If a significant fraction is ozonolysis, then RH will change the gas-phase product distribution as well.

Response: In the original version of the manuscript, we have already assessed and discussed the fraction of α -pinene that reacted with OH and ozone (see section 2.4 and table 1), based on the mixing ratios of these two oxidants. We have estimated that a great part of α -pinene reacted with OH (on average 0.78 ± 0.07) and noted that the further processing of the first generation products – which don't contain C=C bonds – would almost exclusively proceed through OH oxidation. Accordingly, we concluded that SOA compounds detected under our conditions are mainly from OH chemistry.

We note that the fraction of α -pinene that reacted with OH under low RH (0.79 ± 0.07) and high RH (0.77 ± 0.07) are not statistically different (t-test, $p=0.71$), within our experimental variability. Therefore, we do not expect that differences observed between experiments at low and high RH are due to a change in the prevalent oxidant. We recognize that water vapor may change the oxidation product distributions, via its reaction with the stabilized Criegee intermediates, produced upon the ozonolysis of α -pinene. However, we note that not only the fraction of α -pinene that reacts with O_3 is not substantial (in comparison with that reacted with OH), but also a major fraction of α -pinene Criegees undergoes unimolecular decomposition to form OH and does not react with water (Atkinson and Arey, 2003). Therefore, it is unlikely that a change in RH would sizably modify the distribution of the products formed via gas-phase chemistry.

By contrast, the fraction of α -pinene that reacted with OH is found to be sensitive to the NO_x concentration (as already mentioned in the manuscript). The production of O_3 was faster under high NO_x compared to low NO_x , due to an efficient VOC- NO_x catalytic cycle and a higher fraction of OH was scavenged by NO_2 . Consequently, the fraction of α -pinene that reacted with OH under high NO_x (0.75 ± 0.06) is lower than that under low NO_x (0.83 ± 0.04). These small, but statistically significant (0.08; t-test, $p=0.004$) differences in the contribution of ozone/OH to α -pinene oxidation are expected to explain a part of the differences in SOA yields and chemical composition at low and high NO_x , with a higher fraction of ozonolysis products under high NO_x conditions.

While some of the above information was discussed in the old version of the manuscript, we have modified the new text by adding a new section (3.4) related to the above discussion and modified the old section 2.4, accordingly:

3.4 Prevalent oxidation reagent and its influence on SOA yields and chemical composition

Based on the mixing ratios of OH and ozone, we estimate that a greater part of α -pinene has reacted with OH (on average 0.78 ± 0.07). Moreover, it is worthwhile to mention that the further processing of the first generation products – which do not contain C=C bonds – would almost exclusively proceed through OH oxidation. Accordingly, we conclude that SOA compounds detected are mainly from OH chemistry, independent of the NO_x level and relative humidity.

We note that the fraction of α -pinene that reacted with OH under low RH (0.79 ± 0.07) and high RH (0.77 ± 0.07) are not statistically different (t-test, $p=0.71$), within our experimental

variability. Therefore, we do not expect that differences in SOA yields observed between experiments at low and high RH to be due to a change in the prevalent oxidant. We recognize that water vapor may change the oxidation product distributions, via its reaction with the stabilized Criegee intermediates, produced upon the ozonolysis of α -pinene. However, we note that not only the fraction of α -pinene that reacts with O_3 is not substantial (in comparison with that reacted with OH), but also a major fraction of α -pinene Criegee intermediates undergoes unimolecular decomposition to form OH and does not react with water (Atkinson and Arey, 2003). Therefore, it is unlikely that a change in RH would sizably modify the distribution of the products formed via gas-phase chemistry.

By contrast, the fraction of α -pinene that reacted with OH is found to be sensitive to the NO_x concentration. The production of O_3 was faster under high NO_x (average $[O_3] = 35$ ppb) compared to low NO_x (average $[O_3] = 22$ ppb), due to an efficient VOC- NO_x catalytic cycle. Consequently, the fraction of α -pinene that reacted with OH under high NO_x (0.75 ± 0.06) is lower than that under low NO_x (0.83 ± 0.04). These small, but statistically significant differences (8 percentage points; t-test, $p=0.004$) in the contribution of ozone/OH to α -pinene oxidation are expected to explain a small part of the differences in SOA yields and chemical composition observed at low and high NO_x , with higher fraction of ozonolysis products under high NO_x conditions. Despite this, we expect the influence of NO_x on the fate of RO_2 to be the main driver behind the observed differences in SOA yields and chemical composition between low and high NO_x conditions (because ozonolysis products are only a minor fraction and differences between the two conditions are rather small).

Pg 14, ln 13: It would also be beneficial if the authors can talk about these products in terms of RO_2 reactions. Also, why cite a 2011 modeling study when talking about hydroperoxide and carbonyl products from $RO_2 + HO_2$ and $RO_2 + NO$ chemistry, when the mechanisms were deduced much earlier by Atkinson and many others and are now textbook knowledge?

Response: As we have mentioned above, model compounds are only surrogates. While we recognize that oxidation conditions, e.g. NO_x concentrations, may significantly alter the product distribution, we did not select different sets of these surrogates for the different conditions. This is because:

(1) such separation would implicitly suggest that the chemical composition of the few compounds reported at different conditions can be extrapolated to the bulk OA under our conditions;

(2) such separation would significantly limit the number of surrogates at each condition, increasing the sensitivity of the model to the compounds' selection; and

(3) the model is less sensitive to compounds' chemical structure than to their elemental composition (see below and Li et al. (2016)), e.g. number of oxygen and carbon, which has been taken into account by including fragmentation products.

Instead, the same compounds reported at low and high NO_x are used together with fragmentation products as model inputs for all experiments. The contribution of these products is optimized in the model, so that the modelled SOA yields and O:C ratios at a given RH matches the observations. Therefore, while explicit molecular structures are needed as model inputs, the contribution of these molecules in the model is not expected to reflect their true contribution to SOA at different RH and oxidation conditions, especially based on the way the model is setup. Accordingly, we think that discussing differences in the model compounds' contribution at different conditions would not be suitable.

Based on the reviewer comment, we have replaced the citation of a modelling study, Valorso et al. (2011), by the earlier studies reviewed in (Atkinson, 2000), when discussing the influence of NO_x on SOA bulk composition (page 14 line 13).

Pg. 16, ln 17-20: The authors highlighted the importance of solubility in understanding the RH-dependent SOA yields, but the parameterizations only include volatility. The authors say later on that it's assumed that the hydrophilicity is proportional to the volatility, but that would mean treating a chemical process just like a physical process. Given that the model does not consider aqueous reactions, and how important these reactions have been shown to before SOA (i.e., works of McNeill, Ervens, Carlton, and others), it's not clear to this reviewer that the augmented cases with fragmentation and lower volatility products (i.e., more volatility-driven solutions) give the right answers for the right reasons. How do the authors believe the modeling results would change if solubility and aqueous reactions were directly considered?

Response: The model takes hydrophilicity and solubility into account through the activity coefficients of the partitioning compounds. Condensed phase reactions are not considered, nor the possibility that the reactivity depends on relative humidity. We use the thermodynamic simulations to find out whether partitioning alone can account for the increased yield at high RH. Our results show that at low NO_x conditions, equilibrium partitioning between the gas

and liquid phases can explain most of the increase in SOA yields at high RH. In contrast, at high NO_x , equilibrium partitioning alone could not explain the strong increase in the yields with increased RH. Therefore we conclude that under these conditions additional processes including the reactive uptake of semi-volatile species in the particulate aqueous phase would need to occur to explain the enhanced yield at higher RH.

Table S4: Which chemical (i.e., NO, HO₂) regime do these compounds belong to? Can the authors list the abundances that they derived, for each “NO_x” regime? What were the hydrophilicity parameters that the authors assigned for these compounds?

Response: We cannot assign hydrophilicity parameters to model compounds, because the compounds’ activity coefficients, which determine their solubility and hydrophilicity, depend on the composition of the condensed phases and therefore are obtained iteratively during the phase equilibration calculation. As mentioned above, while explicit molecular structures are needed as model inputs, the way the model is set up does not allow us to discuss the influence of the chemical regime on the compounds distributions.

Anonymous Referee #2

This manuscript presents experimental and modelling efforts in order to describe the formation of secondary organic aerosol (SOA) from photo-oxidation of alpha-pinene under various humidities and NO_x concentrations. Also the role of seed aerosol composition is studied in relation to liquid water content of particles. The experiments clearly show that varying the above parameters cause large, and complex, changes in the SOA yields. The authors then attempt to draw further conclusions by phase partitioning calculations. While I find the model results to be less convincing, the model is described in detail, and thus readers can assess the validity of the different assumptions adequately. The paper fits the scope of ACP, and should be considered for publication following the below comments.

General comments:

While the language of the manuscript is very good, I found several descriptions and conclusions hard to follow. Especially the causality in certain sentences should be clarified. As an example, in the abstract it is stated “At low NO_x conditions, equilibrium partitioning between the gas and liquid phases can explain most of the increase in SOA yields at high RH. This is indicated by the model results, when in addition to the α -pinene photooxidation products described in the literature, more fragmented and oxidized organic compounds are added to the model mixtures”. Is the point here that if adding oxidized fragments to the mixture (but not otherwise), the model can explain the increased yields at low NO_x by equilibrium partitioning? The formulation of “indicated ... when” is presumably the main reason for my confusion. Another example is page 15, lines 24-27: “Accordingly, additional insights into the prevalent mechanisms by which the compounds form and evolve can be gained. For example, highly oxygenated compounds cannot be very volatile without significant fragmentation, whereas oligomerization leads to a significant decrease in the compounds’ vapor pressure without necessarily increasing their O:C ratios.” It is unclear to me how the latter sentence is an insight gained from this work? And the content is in any case quite common knowledge, to some extent even used as an assumption in this work. There are several paragraphs with similar issues in the paper, and I recommend the authors (or preferably even someone external) read through the paper with the aim to check how claims of causality are presented.

Title: Currently, the title only reflects the experimental findings, while more focus is put on the model results in the text itself as well as the abstract. Also, the claim is left too general: is this

true regardless of the oxidant (OH, ozone, nitrate radical) or [NO] (only NO_x is mentioned). I suggest revising the title to better describe the content of the paper.

Response: Based on both reviewers' comments, we have significantly modified some parts of the text, especially those related to the model description section. In addition, we have modified the abstract, based on the above comment, as follows:

At low NO_x conditions, equilibrium partitioning between the gas and liquid phases can explain most of the increase in SOA yields observed at high RH, when in addition to the α -pinene photo-oxidation products described in the literature, fragmentation products are added to the model mixtures. This increase is driven by both the increase in the absorptive mass and the solution non-ideality described by the compounds' activity coefficients. In contrast, at high NO_x, equilibrium partitioning alone could not explain the strong increase in the yields with RH. This suggests that other processes, e.g. reactive uptake of semi-volatile species into the liquid phase, may occur and be enhanced at higher RH, especially for compounds formed under high NO_x conditions, e.g. carbonyls.

Based on the comment above, we also modified the sentence on page 15, lines 24-27:

Nevertheless, fitting both organic yields and O:C ratios significantly aid constraining the type of compounds that participate in partitioning (i.e. from a compound O:C ratio and vapor pressure its carbon number can be inferred). For example, highly oxygenated compounds cannot be very volatile without significant fragmentation, whereas oligomerization leads to a significant decrease in the compounds' vapor pressure without necessarily increasing their O:C ratios.

We have also modified the title as follows:

Assessing the influence of NO_x concentrations and relative humidity on secondary organic aerosol yields from α -pinene photo-oxidation through smog chamber experiments and modelling calculations

In the text we have also discussed more thoroughly the influence of the oxidants and NO_x on our conclusions (see above).

Specific comments:

*P1, line 20: For terminology: SOA yields are not affected by particle wall loss. The measured SOA mass is, and if not accounting for that, one will get an *apparent* yield that is too low.*

On the other hand, vapor losses will affect the SOA yields in much more complicated ways. While I do not expect the authors to include vapor wall losses into the model at this stage, and it would be extremely hard to do correctly, the authors should acknowledge that there is a wealth of evidence from the last years that neglecting vapor wall loss will influence SOA yields, including e.g. Kokkola et al., 2014 (the first in a line of recent publications on the role of walls in Teflon chambers), Ehn et al., 2014 (detection of “ELVOC” that irreversibly are lost to walls) and Krechmer et al., 2016 (direct measurements of vapor wall losses in a Teflon chamber). The authors should at least note some of these papers and their findings in the manuscript, rather than only citing the papers that support their approach.

Response: Based on the reviewer comment the sentence on P1, line 20, has been modified in the corrected version of the manuscript as follows:

We used a Monte-Carlo approach to parameterize smog chamber SOA yields as a function of the condensed phase absorptive mass, which includes the sum of OA and the corresponding bound liquid water content.

Also, we have already cited four papers related to wall-losses, but we will add the references suggested by both reviewers. We have significantly changed the section related to the vapour wall losses, where these additional citations were added. The new text reads as follows:

Eq. (8) does not take into consideration the loss of SOA-forming vapours onto the clean Teflon walls, which may suppress SOA yields from laboratory chambers under certain conditions. These processes may be related to the vapours’ reactive uptake onto walls (Loza et al., 2010; Nguyen et al., 2016) or to their absorptive uptake (Zhang et al., 2014; Nah et al., 2016). The reactive uptake of organic vapours onto chamber walls is only significant at high RH. Loss rates for important reactive gases, including glyoxal, epoxides and peroxides have been documented at different RH (Loza et al., 2010; Nguyen et al., 2016). While these processes may also influence reactive SOA-forming compounds under our conditions, they occur at time-scales of hours (Nguyen et al., 2016), much longer compared to the time scales of the absorptive uptake, e.g. based on recent direct measurements of vapour losses onto Teflon walls (~10 min (Krechmer et al., 2016)).

The absorption of organic compounds onto the chamber walls obeys Henry’s law and depends on the compounds’ accommodation coefficients and their activity at the wall/gas interface (see for example Zhang et al., 2015). The dependence of the compounds’ absorption on RH (due to a change in accommodation coefficients or in the activity of the wall absorbed compounds) has not yet been reported to the best of our knowledge and indeed merits further

investigations that are beyond the scope of the current study. Nevertheless, we believe that vapour absorption onto the walls is unlikely to be significantly affected by RH, due to the hydrophobic nature of Teflon and its minor interaction with water under subsaturation conditions ($RH < 80\%$). In our case, we have maintained chamber conditions during our experiments such that vapour wall losses and their inter-experimental differences can be minimized as much as possible. This is done by (1) maintaining a relatively constant wall-to-seed surface ratio for all experiments to avoid systematic biases between experiments and (2) increasing SOA production rates, which rapidly provide a significant particle condensational sink into which condensable gases can partition. We also note that vapour wall losses were found to be minor for the α -pinene SOA system where SOA formation is dominated by quasi-equilibrium growth (Zhang et al., 2014; Nah et al., 2016).

P1, line 20: “as a function of absorptive masses combining organics and the bound liquid water content.” This is a confusing formulation. Rather say “...absorptive mass, defined as the sum of organics and the ...”.

Response: This has been modified in the revised version of the manuscript as follows:

We used a Monte-Carlo approach to parameterize smog chamber SOA yields as a function of the condensed phase absorptive mass, which includes the sum of OA and the corresponding bound liquid water content.

P2, 18: Why limit this statement to semi-volatile species?

Response: In the revised version of the manuscript we remove “semi-volatile”

P5, 5-7: What does it mean when stating “similar NO_x/VOC” when no α -pinene is added??

Response: NO_x/VOC was replaced by NO_x.

P7, 32-33: What were the ozone concentrations? I would like to see a (supplementary) figure with an example experiment showing at least α -pinene, ozone, butanol and OH concentrations together with the SOA mass. Fig. 5: Why are figures 5 and 6 discussed before figures 3 and 4?

Response: O₃ average concentrations were 22 and 35ppb at low NO_x and high NO_x, respectively. In the revised version of the manuscript we have modified section 2.4 (P7, lines 32-33), and added an entire new section, section 3.4, where we discuss the relative importance of O₃ and OH for the oxidation of α -pinene at different conditions:

3.4 Prevalent oxidation reagent and its influence on SOA yields and chemical composition

Based on the mixing ratios of OH and ozone, we estimate that a greater part of α -pinene has reacted with OH (on average 0.78 ± 0.07). Moreover, it is worthwhile to mention that the further processing of the first generation products – which do not contain C=C bonds – would almost exclusively proceed through OH oxidation. Accordingly, we conclude that SOA compounds detected are mainly from OH chemistry, independent of the NO_x level and relative humidity.

We note that the fraction of α -pinene that reacted with OH under low RH (0.79 ± 0.07) and high RH (0.77 ± 0.07) are not statistically different (t-test, $p=0.71$), within our experimental variability. Therefore, we do not expect that differences in SOA yields observed between experiments at low and high RH to be due to a change in the prevalent oxidant. We recognize that water vapor may change the oxidation product distributions, via its reaction with the stabilized Criegee intermediates, produced upon the ozonolysis of α -pinene. However, we note that not only the fraction of α -pinene that reacts with O₃ is not substantial (in comparison with that reacted with OH), but also a major fraction of α -pinene Criegee intermediates undergoes unimolecular decomposition to form OH and does not react with water (Atkinson and Arey, 2003). Therefore, it is unlikely that a change in RH would sizably modify the distribution of the products formed via gas-phase chemistry.

By contrast, the fraction of α -pinene that reacted with OH is found to be sensitive to the NO_x concentration. The production of O₃ was faster under high NO_x (average [O₃] = 35 ppb) compared to low NO_x (average [O₃] = 22 ppb), due to an efficient VOC-NO_x catalytic cycle. Consequently, the fraction of α -pinene that reacted with OH under high NO_x (0.75 ± 0.06) is lower than that under low NO_x (0.83 ± 0.04). These small, but statistically significant differences (8 percentage points; t-test, $p=0.004$) in the contribution of ozone/OH to α -pinene oxidation are expected to explain a small part of the differences in SOA yields and chemical composition observed at low and high NO_x, with higher fraction of ozonolysis products under high NO_x conditions. Despite this, we expect the influence of NO_x on the fate of RO₂ to be the main driver behind the observed differences in SOA yields and chemical composition

between low and high NO_x conditions (because ozonolysis products are only a minor fraction and differences between the two conditions are rather small).

P12, 9: Please use another word than “corresponding” for these comparisons. It is misleading.

Response: The text was modified as follows:

RH was then changed in the model and the effects of RH on SOA yields and degree of oxygenation were evaluated.

P13, 1-10: This is also consistent with more α -pinene producing more SOA and thereby condensation sink (CS), which in turn can more efficiently compete with the walls as a sink of low-volatile vapors. See e.g. the papers cited above. The authors can easily do a sensitivity check for this effect. If there is a large change in particle vs wall loss rates, point 1 should be reconsidered. On the other hand, if the initial seed provides a relatively constant CS (compared to wall losses) for every experiment, then the claim in point 2 that seed concentrations in general are not important seems unjustified, since this was not probed at all in these experiments.

Response: The conclusions in point 2 are based on the fact that we did not observe a correlation between the yields and the initial seed mass or surface concentration, which span a factor of 2.5 (e.g. higher than the span of α -pinene concentration, see point 1). Therefore, would there be an influence of the seed surface we would have observed it, but this was not the case. As mentioned in the manuscript, this is consistent with new results suggesting that for the α -pinene system SOA formation is dominated by quasi-equilibrium growth and vapour losses to the walls do not depend on the seed concentrations, but rather on SOA formation rates (Nah et al., 2016).

In Point 1, we simply describe first the observation that SOA yields increase with α -pinene concentrations. This is expected from the equilibrium growth of α -pinene SOA from condensable semi-volatile vapours, and is consistent with point 2 (this is the basis of the partitioning theory). We have now assessed the change in the particle condensation sink between experiments with α -pinene concentrations equal 20 and 30ppb. The change is too negligible, ~10%, to have an impact on the ratio between the particle and wall surfaces and on the vapour wall losses. Therefore, we think that point 1 is still valid.

P13, 2: Please clearly distinguish between percentages and percentage points. I expect this should be the latter.

Response: It is percentage point, as stated by the reviewer. In the corrected version of the manuscript we add “(percentage point)” after the 2%.

Fig. 4: There is a clear bimodal distribution for most cases, which is also noted in the manuscript. However, I find the explanations and discussion about it lacking. The authors state that particle number increased, but no new particle formation was observed. This needs some further discussion. Where do the particles come from then? Additionally, the bimodality is used as proof for LLPS, which the model also predicts, but I do not find a clear description of why the organics form this bimodal distribution. Fig. 7&8: Make the contrast between the currently light and full colors more visible. At least on my screen some pairs were hard to distinguish.

Response: The development of a bimodal distribution is described in section 3.3, pages 14 – 15 and then discussed on page 20, lines 14 – 26. The aerosol size-resolved chemical composition is used only to infer the behaviour of the absorptive organic phase and its mixing with the inorganic seed, rather than to model the particle dynamics in the chamber and the processes involved (in our opinion the latter needs much more constraints, which we do not have). As we have mentioned (on page 15, line 2), we have observed an increase in the total particle number, but based on Fig. S14, we did not observe clear and intense nucleation events (substantial increase in the number of small particles). Therefore, while we cannot exclude the occurrence of moderate new particle formation events, we believe that we do not have the adequate tools (detection of new particles below 10 nm) to confirm it. We also cannot exclude that the transmission efficiency of smaller particles in the lines increases upon their growth, which would also explain the increase in the particle number. Accordingly, we prefer keeping the same statements in the text, without explicitly affirming that new particle events occurred, especially that whether nucleation occurred or not does not change in our opinion the conclusions of the paper.

For all experiments, the aerosol size distributions show two externally mixed aerosol populations, with a mode at lower diameters (~200 nm, mode 1) mostly containing SOA and another mode at higher diameters (~400 nm, mode 2) mostly consisting of the seed. As

mentioned on page 20, the formation of these two populations may occur by the homogeneous or heterogeneous nucleation of highly oxidized non-volatile products. Homogeneous nucleation implies new particle formation (which we did not exclude as mentioned above), while heterogeneous nucleation proceeds via condensational growth of non-volatile vapours onto smaller particles (the particle surface mode). Both processes are expected to create small organic rich particles, providing an organic absorptive phase into which additional semi-volatile compounds may preferentially partition. While we cannot exclude any of the two processes (homogeneous or heterogeneous nucleation), the size-resolved chemical composition information is used here to provide compelling evidence that the organic and the electrolyte phases are not mixed, confirming the modelling results.

P15, 36-37 and related O:C discussion: A variation of 0.03 from 0.45 to 0.48 is considered “almost constant” while an increase of 0.08 is considered significant? And only several pages later in section 4.6 is it noted that the uncertainty in O:C is 20-30%. It is also noted that the O:C values are likely biased low since the latest parametrization for O:C calculations are not used. These things should be mentioned earlier, so a reader can properly assess the meaningfulness of the comparisons done in sections 4.2-4.3. Considering all the above, the tuning of the model to match these values does not in my mind give much more insight into the formation mechanisms of SOA in this system.

Response: In general, while it is hard to assess the uncertainties related to the AMS O:C measurements, a distinction should be made between measurement precision and accuracy. We do not expect the accuracy of the O:C ratios determined by the HR-ToF-AMS to be less than 20%, while changes in the O:C ratios can be detected more precisely by the instrument (~1-2%). As noted by the reviewer, we have used an earlier parameterization to determine the O:C ratio (Aiken et al., 2008), and mentioned that the later parameterization would yield higher O:C values (Canagaratna et al., 2015). While it is not straightforward (impossible) to judge which parameterization is better for our system, based on our recent measurements of ambient samples (see Bozzetti et al. (2017)), the later parameterization does not seem to provide more accurate results. Therefore, we decided to maintain the earlier parameterization.

As suggested by the reviewer, in the corrected version of the manuscript we will discuss earlier in the text the uncertainties related to the determination of the O:C ratios. In the method section (2.1), the following text is added:

The HR-ToF-AMS data were processed and analysed using the analysis software SQUIRREL (SeQUential Igor data RetRiEvaL) v.1.52L and PIKA (Peak Integration by Key Analysis) v.1.11L for IGOR Pro software package (Wavemetrics, Inc., Portland, OR, USA). From the HR analysis of the mass spectra, the O:C ratios of the bulk OA were determined based on the parameterization proposed by Aiken et al. (2008). We note that while the assessment of the uncertainties related to the O:C measurements by the HR-ToF-AMS is not straightforward, a distinction should be made between measurement precision and accuracy. We do not expect the accuracy of the O:C ratios determined by the HR-ToF-AMS to be less than ~20 % (Aiken et al., 2008; Pieber et al., 2016; Canagaratna et al., 2015; Bozzetti et al., 2017). For example, the use of a more recent parameterization (Canagaratna et al., 2015) would yield higher O:C values (by 18 %) and the O:C ratios reported here may be regarded as lowest estimates. By contrast, relative changes in the O:C ratios are expected to be detected more precisely by the instrument (~1-2 %). The influence of potential biases and uncertainties in the determination of the O:C ratios on our results will be discussed in the text.

The first statement mentioned by the reviewer reads as follows: “This agreement was achieved although the model could not reproduce the high values of nor the change in the O:C ratios with RH (the modelled O:C remained almost constant at 0.45-0.48 at low and high RH, while the measured O:C increased from 0.56 at low RH to 0.64 at high RH).” The point here is that the model underestimated the O:C values by 24%, and predicted a smaller increase in their values with RH (6% compared to 14%). This would be even worse if the more recent parameterization of Canagaratna et al. (2015) was used. The above sentence is modified as follows:

This agreement was achieved although the model underestimated the O:C ratios by 24 % and predicted a smaller increase in the O:C values with RH than observed. The modelled O:C remained almost constant at 0.45-0.48 at low and high RH (difference of 6 %), while the measured O:C increased from 0.56 at low RH to 0.64 at high RH (change of 14 %). We also note that the use of the more recent parameterization by Canagaratna et al. (2015) would yield even higher O:C values, widening the gap between measured and modelled O:C ratios.

The other point raised by the reviewer pertains to the discussion of the differences between the two parameterizations on page 19. The point here is that the use of the most recent parameterization would yield higher O:C ratios, which would require including more fragmented products in the model than already done to reproduce the measured O:C ratios. Therefore, using either parameterization will not change the conclusions of the paper in a

qualitative term, suggesting that fragmentation products should be included in the model to reproduce the observations. In the discussion section, we added the following clarifications:

The measured O:C ratio is a key parameter for constraining the model. Here, we have used the high resolution parameterization proposed by Aiken et al. (2007), while the use of the more recent parameterization by Canagaratna et al. (2015) would result in even higher O:C ratios (by 18 %). Higher O:C ratios would require increasing even further the contribution or the degree of oxidation of the fragmented compounds and would imply that the model predicts an even higher sensitivity of the yields to RH. Therefore, the O:C values used here yield more conservative estimates of the contributions (or the degree of oxidation) of fragmented products in the model and of the sensitivity of the yields and O:C ratios to the RH.

P18, 3-5: Is this shown somewhere, or just stated? There were also other places where the formulations are such that I expect there to be a figure showing the result. The authors should consider adding “not shown” in places where the information is not visible in any plot or table.

Response: This is indeed stated and not shown. In the revised version of the manuscript we have added “(not shown)” here and in places where the information is not visible in any plot or table.

P18, 22-26. There are too many numbers listed in the text, and this is especially true here. Please consider rewriting this.

Response: The section reads as follows: “This increased partitioning is limited by an increase in the activity coefficients of these compounds (for experiment 8 from 1.69 and 1.63 at low RH to 2.70 and 2.93 at high RH and for experiment 13 from 1.49 and 1.54 at low RH to 2.88 and 3.24 at high RH). Conversely, the 5-fold enhanced partitioning of the fragmented and more functionalized compounds (5-COOH-3-OH-pentanal and succinic acid) into the condensed phase at high RH is driven by the increase of the absorptive mass and the slight decrease of the compounds’ activity coefficients (for experiment 8 from 0.84 and 0.51 at low RH to 0.68 and 0.43 at high RH and for experiment 13 from 0.90 and 0.56 at low RH to 0.69 and 0.44 at high RH).” We agree that the text contains a lot of numbers, but all of them are between brackets, so that the reader is not obliged to read them. We think these numbers are

necessary to provide an idea about the range of the modelled activity coefficients, under different RH and for different types of compounds.

P18, 19-22 and Fig. 9: It is hard to follow discussion about increase of factors 2-5 from a plot ranging 14 orders of magnitude. Could Fig. 9 be moved to the SI, and some more specific plot included in the main text?

Response: Figure 9 shows the partitioning of the individual model compounds to the different phases, which provides an idea which compounds are expected to contribute to the two particle phases under different conditions. We think this figure is required in the main text.

P18, 39-40: I do not understand at all what this sentence is supposed to say.

Response: This sentence has been removed from the manuscript.

P19, 7: “sufficient” for what?

Response: We specify “sufficient to represent the species within one volatility bin”.

P20, 22: Expected based on what?

Response: This section has been reformulated as follows:

Considering the size-resolved particle chemical composition discussed in Sect. 3.3 (Figure 4), LLPS is likely not realized within single particles but the aerosol population splits up into a predominantly organic mode at ~200 nm and a predominantly inorganic mode at ~400 nm. The formation of these two populations may occur by the homogeneous or heterogeneous nucleation of highly oxidized non-volatile products. Homogeneous nucleation implies new particle formation (which would only be moderate – see Figures S13 and S14 – due to the high condensation sink in the chamber), while heterogeneous nucleation proceeds via condensational growth (which would occur on smaller particles with higher surface). Both processes are expected to create small organic rich particles, providing an organic absorptive phase into which additional semi-volatile compounds may preferentially partition. When the organic and electrolyte phases are present in different particles the two phases communicate via gas phase diffusion, and equilibration time-scales depend on the components’ volatility.

For compounds with $C^* = 0.1-100 \mu\text{g m}^{-3}$ equilibration occurs within time-scales of minutes to tens of minutes, assuming no bulk phase diffusion limitations (Marcolli et al., 2004). In the larger particle electrolyte-rich mode, the inorganic ions would exert a salting-out effect driving the organic compounds to partition to the gas-phase or into the smaller organic-rich particles. This would prevent the organic compounds from partitioning in significant amounts into the seed aerosol from the beginning and deplete even further these larger particles from the organic material. Under such a scenario an externally-mixed phase-separated aerosol may evolve in the smog chamber. Overall, the size-resolved chemical composition information confirms the modelling results providing compelling evidence for organic-electrolyte LLPS.

P21, 28-29: This is too strong a statement in my opinion. Rather say that only with inclusion of the fragments could your model describe both SOA mass and O:C.

Response: This has been modified in the revised version of the manuscript as suggested by the reviewer:

Our results show that only with the inclusion of fragmentation products the model could simultaneously explain SOA concentrations and O:C ratios.

P21, 40-P22, 2: Such a statement should be included much earlier in the discussions on bimodality, and not saved to the last lines of the manuscript.

Response: This statement has been already discussed earlier in the manuscript on page 20 and modified in the new version of the text based on an earlier comment of the reviewer. The section on page 20 now reads as follows:

Considering the size-resolved particle chemical composition discussed in Sect. 3.3 (Figure 4), LLPS is likely not realized within single particles but the aerosol population splits up into a predominantly organic mode at ~200 nm and a predominantly inorganic mode at ~400 nm. The formation of these two populations may occur by the homogeneous or heterogeneous nucleation of highly oxidized non-volatile products. Homogeneous nucleation implies new particle formation (which would only be moderate – see Figures S13 and S14 – due to the high condensation sink in the chamber), while heterogeneous nucleation proceeds via condensational growth (which would occur on smaller particles with higher surface). Both processes are expected to create small organic rich particles, providing an organic absorptive phase into which additional semi-volatile compounds may preferentially partition. When the

organic and electrolyte phases are present in different particles the two phases communicate via gas phase diffusion, and equilibration time-scales depend on the components' volatility. For compounds with $C^* = 0.1-100 \mu\text{g m}^{-3}$ equilibration occurs within time-scales of minutes to tens of minutes, assuming no bulk phase diffusion limitations (Marcolli et al., 2004). In the larger particle electrolyte-rich mode, the inorganic ions would exert a salting-out effect driving the organic compounds to partition to the gas-phase or into the smaller organic-rich particles. This would prevent the organic compounds from partitioning in significant amounts into the seed aerosol from the beginning and deplete even further these larger particles from the organic material. Under such a scenario an externally-mixed phase-separated aerosol may evolve in the smog chamber. Overall, the size-resolved chemical composition information confirms the modelling results providing compelling evidence for organic-electrolyte LLPS.

P22, 2: Again, what is this expectation based on? Work in this paper or other work?

Response: This is based on our work and has been clarified in the new version of the manuscript.

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Assessing the influence of NO_x concentrations and relative humidity on α -Pinene secondary organic aerosol yields from α -pinene photo-oxidation through smog chamber experiments and modelling calculations increase at higher relative humidity and low NO_x conditions

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Abstract. Secondary organic aerosol (SOA) yields from the photooxidation of α -pinene were investigated in smog chamber (SC) experiments at low (23-29 %) and high (60-69 %) relative humidity (RH), various
20 NO_x/VOC ratios (0.04-3.8) and with different aerosol seed chemical compositions (acidic to neutralized sulfate-containing or hydrophobic organic). A combination of a scanning mobility particle sizer and an Aerodyne high resolution time-of-flight aerosol mass spectrometer was used to determine SOA mass concentration and chemical composition. ~~We used a Monte-Carlo approach to parameterize smog chamber SOA~~ ~~We present wall-loss-corrected~~ yields as a function of the condensed phase absorptive mass, which includes the sum of OA masses
25 ~~combining organics~~ and the corresponding bound liquid water content. High RH increased SOA yields by up to six times (1.5-6.4) compared to low RH. The yields at low NO_x/VOC ratios were in general higher compared to yields at high NO_x/VOC ratios. This NO_x dependence follows the same trend as seen in previous studies for α -pinene SOA.

A novel approach of data evaluation using volatility distributions derived from experimental data served as basis
30 for thermodynamic phase partitioning calculations of model mixtures in this study. These calculations predict liquid-liquid phase separation into organic-rich and electrolyte phases. At low NO_x conditions, equilibrium partitioning between the gas and liquid phases can explain most of the increase in SOA yields observed at high RH. ~~This is indicated by the model results~~, when in addition to the α -pinene photo-oxidation ~~photooxidation~~

products described in the literature, ~~fragmentation products more fragmented and oxidized organic compounds~~ are added to the model mixtures. This increase is driven by both the increase in the absorptive mass ~~due to the additional particulate water~~ and the solution non-ideality described by the ~~compounds'~~ activity coefficients. In contrast, at high NO_x, equilibrium partitioning alone could not explain the strong increase in the yields with ~~increased~~ RH. This suggests that other processes, ~~e.g. including the~~ reactive uptake of semi-volatile species into the liquid phase, may occur and be enhanced at higher RH, especially for compounds formed under high NO_x conditions, ~~e.g. such as~~ carbonyls.

1 Introduction

Organic aerosol (OA) accounts for 20 to 90 % of the submicron ambient aerosol (Jimenez et al., 2009 and references therein), a great part of which is secondary organic aerosol (SOA) formed via the condensation of oxidation products of gas-phase precursors. Several direct (e.g. radiocarbon dating) and indirect observations underline the key role of biogenic volatile organic compounds (VOC) for SOA formation (El Haddad et al., 2013 and references therein). Current state-of-the-art models are unable to predict the burden of biogenic SOA, especially in urban atmospheres (Hoyle et al., 2011), highlighting a fundamental deficit in our knowledge of the chemical pathways by which SOA accumulates and evolves in the atmosphere.

The ensemble of gaseous and particulate phase species involved in SOA formation is immensely complex. The compounds relevant for SOA formation are often a minor fraction, resulting in yields (SOA formed to precursor reacted) of only a few percent. Their chemical composition and volatility distribution strongly depend on the oxidation conditions, most notably on the fate of organic peroxy radicals (RO₂), which either react with nitrogen oxides (NO_x) or other peroxy radicals (RO₂ and HO₂). The influence of NO_x on the oxidation mechanisms and SOA formation is commonly described by the NO_x/VOC ratio and has been under close scrutiny lately. For most light precursors, such as isoprene and monoterpenes (including α -pinene), SOA yields appear to be strongly influenced by the NO_x/VOC ratio, with a general enhancement observed under low NO_x conditions (Ng et al., 2007; Presto et al., 2005).

~~Species~~ ~~Semi-volatile species~~ involved in SOA formation are subject to ongoing chemical degradation, which may lead to compounds of either lower (when functionalization dominates) or higher volatility (when fragmentation dominates). As a consequence, SOA yields and degrees of oxygenation (described by the atomic oxygen to carbon ratio O:C) may depend on the extent to which these species were exposed to oxidants (Kroll and Seinfeld, 2008; Donahue et al., 2012a).

SOA yields are generally described by the absorptive equilibrium partitioning of condensable species to a well-mixed liquid phase (Odum et al., 1996), which depends upon the chemical species' saturation vapor pressures (classically two model products are considered) and their liquid-phase activities (modified Raoult's law). Donahue and co-workers proposed the use of a "volatility basis set" (VBS) for a better representation of the wide range of OA in the atmosphere and the ongoing oxidation of semi-volatile organics (Donahue et al., 2011; 2012b and references therein). However, several difficulties remain in estimating or measuring the saturation vapor pressures (e.g. Huisman et al., 2013; Bilde et al., 2015), activity coefficients and the mean molecular weight of the condensing species (e.g. Clegg et al., 2008a, b).

Difficulties increase when considering the role of relative humidity (RH) and of electrolyte particles on organic partitioning (Zuend and Seinfeld, 2013 and references therein). From a thermodynamic point of view, water interacts with SOA components by altering the water content of aerosol particles (also at subsaturated conditions) and hence the equilibrium concentration of water-soluble organic compounds. According to the equilibrium equation of Pankow (1994), an increase in the condensed fraction of the organics can be achieved by 1) increasing the absorptive particulate mass, 2) decreasing the average molecular weight of condensed species or 3) decreasing the activity coefficients of organic species (Pankow, 1994). Therefore, it is expected that an increase in particulate water content would in principle enhance SOA yields of water-miscible species. Model calculations predict a pronounced effect, especially at low organic mass loadings (Pankow, 2010). However, literature reports based on experimental results seem contradictory: RH dependent yields have been reported for the ozonolysis of limonene, α -pinene and Δ^3 -carene (Jonsson et al., 2006), while Prisle et al. (2010) found a negligible impact of RH on SOA yields from α -pinene ozonolysis. A substantial effect of RH on SOA yields was observed only for studies at precursor concentrations <1000 ppbv and with large variation in RH (0.01 % & 31 % RH (Bonn et al., 2002); < 2–58 % RH (Cocker et al., 2001); < 2–85 % RH (Jonsson et al., 2006).

The impact of RH on the partitioning of organic species may be severely suppressed by considerable deviations from ideal mixing between the condensing organic species and the prevailing condensed phase. A growing number of studies show that organic compounds are salted out in internally mixed organic/ inorganic/ water aerosol particles and two stable liquid phases may develop: an aqueous electrolyte solution and an organic solution (Marcolli and Krieger, 2006; You et al., 2014). The miscibility of an organic compound in aqueous droplets containing electrolytes depends on numerous factors including temperature, relative humidity, organic compound polarity, relative contribution of the compound to the bulk particulate matter and the chemical nature of the electrolytes (You et al., 2013; Zuend and Seinfeld, 2013). For example, phase separation was observed to always occur for organic compounds with an O:C < 0.5 and at low relative humidity (You et al., 2013), and is especially pronounced in the case of ammonium sulfate (compared to ammonium hydrogen sulfate and nitrate). While neglecting phase separation of organic compounds and the effect of RH thereon bears the potential for invalid yield predictions (e.g. if the condensed phase is considered to comprise a mixed electrolyte and organic solution), the way the complex organic matrix interacts with water and the inorganic species remains virtually unknown.

Another complication that might influence the interpretation of chamber experiments conducted at different RH is the enhancement of SOA yields by the potential reactive uptake of organic products into the particle phase (Kroll and Seinfeld, 2008; Kleindienst et al., 2006; Jang et al., 2002; Iinuma et al., 2007). The mechanisms by which such reactions occur are not fully identified, but may involve an ester or an aldol formation, which are expected to be favourable in the absence of water and are possibly catalytic under acidic conditions. Assessing the relative importance of particle phase processing at different particle water contents would require decoupling SOA thermodynamics and additional reactivity.

In this study, we examine the impact of particle water content on the chemical composition and yields of α -pinene SOA formed under low and high NO_x . This is performed by varying NO_x/α -pinene ratios, aerosol seed composition (hygroscopicity, acidity) and relative humidity. Results are parameterized within a thermodynamic framework to investigate whether changes in SOA non-ideal mixing properties with particle water content may

explain the variation in SOA yields with RH. SOA yields reported here may aid the parameterization of the NO_x and particulate water dependence of α -pinene SOA production for further use in atmospheric models.

2 Methods

2.1 Experimental setup and instrumentation

20 experiments, listed in Table 1, were carried out in the smog chamber (SC) of the Paul Scherrer Institute (PSI): a Teflon bag of 27 m³ suspended in a temperature-controlled housing (Paulsen et al., 2005). Photochemistry was initiated by four xenon arc lamps (4 kW rated power, 1.55·10⁵ lumens each, XBO 4000 W/HS, OSRAM), facing parallel to the SC bag, and emitting a light spectrum similar to the solar spectrum, and 80 black lights (Philips, Cleo performance 100 W) to accelerate the aging process, located underneath the SC bag, with emission between 300–400 nm wavelength (light characterisation in Platt et al. (2013)). A reflecting aluminum foil surrounds the SC bag to maintain light intensity and light diffusion.

10 Various parameters were monitored in the SC. The temperature (T) and RH measurement was optimized by passing SC air through a radiation shielded sensor. One of two different high resolution time-of-flight aerosol mass spectrometers (HR-ToF-AMS, Aerodyne Research, Inc., Billerica, MA, USA) was operated during three different campaigns to measure online size-resolved chemical composition (organics, ammonium, nitrate, sulfate, chloride) of non-refractory particles (DeCarlo et al., 2006). The HR-ToF-AMS were equipped with two different PM_{2.5} lenses (Williams et al., 2013) to sample particles up to large diameters above 1 μm. The sampled aerosol was dried (~10 % RH) before measurement. A supporting flow of ~1.5 L min⁻¹ was maintained parallel to the HR-ToF-AMS to minimize diffusive losses in the sampling lines.

The HR-ToF-AMS data were processed and analysed using the analysis software SQUIRREL (SeQUential Igor data RetRiEval) v.1.52L and PIKA (Peak Integration by Key Analysis) v.1.11L~~was corrected~~ for IGOR Pro software package (Wavemetrics, Inc., Portland, OR, USA). From the HR analysis of the mass spectra, the O:C ratios of the bulk OA were determined based on the parameterization proposed by Aiken et al. . We note that while the assessment of the uncertainties related to the O:C measurements by the HR-ToF-AMS is not straightforward, a distinction should be made between measurement precision and accuracy. We do not expect the accuracy of the O:C ratios determined by the HR-ToF-AMS to be less than ~20 % . For example, the use of a more recent parameterization would yield higher O:C values (by 18 %) and the O:C ratios reported here may be regarded as lowest estimates. By contrast, relative changes in the O:C ratios are expected to be detected more precisely by the instrument (~1-2 %). The influence of potential biases and uncertainties in the determination of the O:C ratios on our results will be discussed in the text~~gas-phase contributions.~~

Two scanning mobility particle sizers (SMPS) were additionally deployed for the measurement of the aerosol size distributions. The first SMPS (a custom built differential mobility analyzer, DMA: extended length $L_{\text{eff}} = 0.93$ cm, $d_{\text{m max}}=1000$ nm, recirculating sheath flow, and a condensation particle counter, CPC 3022 (TSI)) was connected to the HR-ToF-AMS sampling line to analyze the dried particles. A second SMPS (SMPS_{wet}, a TSI DMA classifier 3081 with recirculating sheath flow and a TSI CPC 3022A) and a CPC (TSI: CPC 3025A) measured the wet particle number size distribution and total number concentration ($d > 3$ nm), respectively.

35 Gas-phase compounds with a higher proton affinity than water (166.5 kcal mol⁻¹) were measured with a quadrupole proton transfer reaction mass spectrometer (PTR-MS, Ionicon). The PTR-MS was calibrated before each experiment for α -pinene, detected at m/z 137 and m/z 81; the accuracy of these measurements was estimated to be ~5%, based on the purity indicated on the calibration gas cylinder.

A modified NO_x instrument including a photolytic NO₂-to-NO converter (Thermo Environmental Instruments 42C trace level NO_x analyzer equipped with a blue light converter) and two ozone monitors (Monitor Labs 8810 ozone analyzer, EnviroNics S300 ozone analyzer) monitored NO_x and O₃ in the chamber.

2.2 Chamber operation and aerosol seeding

5 SOA formation and growth from α -pinene was induced by the following SC operation sequence: (1) humidification of the chamber, (2) addition of seed aerosol, (3) introduction of VOCs, (4) addition of nitrous acid (HONO) as an OH precursor, (5) addition of nitrogen oxides (equal amounts of NO + NO₂) if applicable, (6) an equilibration period (30–45 min), (7) switching on of xenon and black lights to generate OH radicals, and (8) a reaction time of 5 to 20 h (corresponding to 0–2 × 10⁸ cm⁻³ h OH exposure, see Sect. 2.4). Experimental
10 conditions for each individual experiment are summarized in Table 1. ~~Prior to~~ ~~Ahead of~~ each experiment, cleaning of the SC was performed by the injection of several ppmv of ozone (5 h) and the simultaneous irradiation with black lights (10 h) at a temperature of 20 °C. This was followed by a pure air flushing period at high relative humidity (~60 %) at a temperature of approximately 30 °C for at least 20 h. Three blank experiments (seed aerosol, lights switched on, high RH, but without adding α -pinene) were carried out to make
15 sure that the organic aerosol formed during the experiments is not significantly influenced by background contamination in the SC. The organic mass concentration formed was substantially lower (< 0.1 up to 2.8 $\mu\text{g m}^{-3}$) than during comparable experiments (similar NO_x, ~~VOC~~ and RH).

In the chamber, the temperature varied between 21 °C and 26 °C. Due to heat from the xenon and black lights, the temperature increased, stabilizing only ~1 h after experiment start. The increase in temperature of 1–4 °C led
20 to an absolute decrease in RH of ~2–20 %, thus the RH range in Table 1, 23–67 %, is given for the ~~temperature-~~ stable ~~temperature-~~ period. α -Pinene (98 %, Aldrich) and an OH reactivity tracer (9-times deuterated butanol, 98 %, D9, Cambridge Isotope Laboratories), hereafter referred to as butanol-d9 (1 μL injected \approx 10 ppbv in the SC), were sequentially injected into an evaporation glass bulb heated to 80 °C. The two VOCs were transferred into the bag by a dilution and flush flow (each 15 L min⁻¹, maintained for 15 min) from an air purifier (737-250
25 series, AADCO Instruments, Inc., USA), further referred to as “pure air”. Initial α -pinene concentrations were 16.1–31.7 ppbv.

HONO was used as a source of both NO and OH, produced by continuous mixing in a reaction vessel of the reagents sodium nitrite (NaNO₂, 1 mmol L⁻¹ in milliQ-H₂O) and three different concentrations of sulfuric acid solutions (H₂SO₄, 1 mmol L⁻¹ (expts. 1–6), 10 mmol L⁻¹ (expts. 7, 8, 11, 13) and 100 mmol L⁻¹ (expts. 9, 10, 12,
30 14) in milliQ-H₂O) (Taira and Kanda, 1990). HONO was carried by 2.5 ± 0.2 L min⁻¹ pure air flow into the SC. 2 ppbv (± 10 %) of HONO were injected before lights on to initiate photochemistry and the injection was continued throughout all experiments. A chemiluminescence-based NO_x instrument (Monitor Labs 9841A NO_x analyzer) was attached to the HONO source to monitor the injected concentration throughout the experiment. In addition, equal concentrations of NO (99.8 %; 1005 ppmv ± 2 %) and NO₂ (purity: 98 %; 1005 ppmv ± 3 %),
35 resulting in 19.6–75.1 ppbv initial NO_x, were added during experiments with NO_x/ α -pinene > 1. Within the results section, the two terms “low NO_x” and “high NO_x” refer to the following conditions:

- ~~Low~~ ~~low~~ NO_x = NO_x/ α -pinene ~~<~~ < 0.1, with continuous HONO injection, indicated by an asterisk in Table 1 and figures.

- ~~High~~ $\text{NO}_x = \text{NO} + \text{NO}_2$ / α -pinene \rightarrow 1: Initial injection of $\text{NO} + \text{NO}_2$ with continuous HONO injection.

In Table 1, we report for high NO_x conditions the initial NO_x concentration, (which decays with time) and for low NO_x conditions the mean NO_x concentration. We note that the NO levels are the main driver for determining whether $\text{RO}_2\text{-NO}$ or $\text{RO}_2\text{-RO}_2$ reactions would prevail. Based on our calculations, RO_2 radicals would predominantly react with NO, when the concentration of the latter is higher than only 1 ppb. These conditions can be considered as high NO_x . During high NO_x experiments and throughout the period when the majority of α -pinene was consumed, the NO concentration remained higher than 5ppb, indicating that α -pinene oxidation proceeded under high NO_x . During low NO_x experiments, the average NO_x concentration was around 1-2 ppb, predominated by NO_2 , while the NO concentrations were below detection limits ($<0.1\text{ppb}$). Under these conditions $\text{RO}_2\text{-RO}_2$ reactions may prevail.

During 14 experiments (no. 1–14) and three blank experiments, an ammonium hydrogen sulfate (NH_4HSO_4 , Aldrich) solution in ultrapure Milli-Q water (1 g L^{-1}) was nebulized (0.6 L min^{-1}) and introduced into the SC with a pure air dilution flow of 10 L min^{-1} to act as seed particles. To keep the seed aerosol in a liquid state, no drier was used behind the nebulizer. We determined the acidity of the seed particles, here described by the ratio NH_4/SO_4 , by the comparison between the HR-ToF-AMS measurements of the seed particles and nebulized $(\text{NH}_4)_2\text{SO}_4$ and NH_4HSO_4 solutions. An NH_4/SO_4 ratio between 1 and 2 indicates a rather neutral seed composition, representing a mixture of NH_4HSO_4 , $(\text{NH}_4)_3(\text{SO}_4)_2$ (letovicite) and $(\text{NH}_4)_2\text{SO}_4$. For simplicity, we replaced $(\text{NH}_4)_3(\text{SO}_4)_2$ by an equal mix of NH_4HSO_4 and $(\text{NH}_4)_2\text{SO}_4$ for the assumption of density and growth factors. By contrast, an $\text{NH}_4/\text{SO}_4 \leq 1$ indicates an acidic seed, consisting of NH_4HSO_4 and H_2SO_4 .

During experiments 7–14, the nebulized NH_4HSO_4 solution was partly neutralized to $(\text{NH}_4)_2\text{SO}_4$, presumably by background NH_3 (Fig. S1, Supplement). During experiments 1–6 the seed was composed of an aqueous mixture of NH_4HSO_4 and H_2SO_4 , due to the in-situ production of gas-phase H_2SO_4 via the HONO injection system, suppressing the seed neutralization by NH_3 (even though particulate H_2SO_4 was minimised by a teflon filter applied between the HONO source and the SC).

Additionally, three α -pinene experiments (no. 15–17) and one blank experiment were conducted using an inert hydrophobic fluorinated hydrocarbon seed (further referred to as CF-seed; $\text{CF}_3\text{CF}_2\text{CF}_2\text{O}[\text{CF}(\text{CF}_3)\text{CF}_2\text{O}]_n\text{CF}_2\text{CF}_3$; Krytox® 1525). The CF-seed was generated via the evaporation of the pure compound at a temperature of $125\text{--}145\text{ }^\circ\text{C}$ and subsequent homogeneous nucleation in a pure air flow of $2.4 \pm 0.1\text{ L min}^{-1}$. The CF-seed concentration was $6.7\text{--}10\text{ }\mu\text{g m}^{-3}$ when lights were switched on, and decayed very rapidly to values below detection limit of the HR-ToF-AMS ($0.1\text{ }\mu\text{g m}^{-3}$) during the course of the experiment. The CF-seed mass spectrum in the HR-ToF-AMS is clearly distinct from that of α -pinene SOA, with main contributions at m/z 69 (CF_3), m/z 169 (C_3F_7) and m/z 119 (C_2F_5) (Fig. S2 and Table S1, Supplement).

Table 1 lists the expected physical state and seed composition of each experiment dependent on RH (AHS = ammonium hydrogen sulfate (NH_4HSO_4); AS = ammonium sulfate ($(\text{NH}_4)_2\text{SO}_4$); SA = sulfuric acid (H_2SO_4); CF = fluorinated carbon. Submicrometer AS particles and letovicite particles ($(\text{NH}_4)_3\text{H}(\text{SO}_4)_2$) are expected to effloresce at about 35 % RH while more acidic particles should remain liquid between 20 and 30 % RH (Martin, 2000; Ciobanu et al., 2010).

2.3 Estimation of the hygroscopic growth factors and liquid water content

The absolute liquid water content (LWC) of the aerosol particles was derived for the bulk aerosol mass and the size resolved mass distribution, based on literature growth factors, the measured RH and chemical composition. The growth factor GF (RH) of a particle is defined as the ratio of the wet diameter at a given RH to the dry diameter: $GF(RH) = d(RH)/d_{dry}$. Inorganic GFs were taken from the Aerosol Diameter Dependent Equilibrium Model (ADDEM, Topping et al., 2005) for diameters of 360 nm. Organic GFs were derived using the relationship between the hygroscopicity parameter κ and the degree of oxygenation [$\kappa = 0.29 \times (O:C)$] from Chang et al. (2010), well representing the hygroscopicity of α -pinene SOA measured by Massoli et al. (2010). The measured degree of oxygenation at an OH exposure of $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$ was used to derive κ , which in turn was converted to GF, assuming a negligible curvature (Kelvin) effect (Kreidenweis et al., 2005):

$$GF(RH) = \left(1 + \kappa \frac{\frac{RH}{100\%}}{1 - \frac{RH}{100\%}} \right)^{\frac{1}{3}} \quad (1)$$

The mixed GFs for aerosol containing inorganic and organic species were determined using as a first approximation the Zdanovskii-Stokes-Robinson (ZSR) volume mixing rule (Stokes and Robinson, 1966):

$$GF_{mixed}(RH) = \left(\sum_i \varepsilon_i \times (GF_i(RH))^3 \right)^{\frac{1}{3}} \quad (2)$$

where ε_i and $GF_i(RH)$ denote the volume fraction and GF (RH) of species i , respectively. The H_2O volume (V_{H_2O}) was calculated using the definition of GF (RH) and the dry volume (V_{dry}):

$$V_{H_2O} = V_{dry} \times (GF(RH))^3 - 1 \quad (3)$$

V_{H_2O} multiplied by the density of water (1 g cm^{-3}) results in the LWC.

The dry (S_{dry}) and wet (S_{wet}) surfaces from HR-ToF-AMS size resolved data were calculated with Eq. (4) and Eq. (5), respectively:

$$S_{dry} = 6 \times (V_{dry})/d_{dry} \quad (4)$$

$$S_{wet} = 6 \times (V_{dry} + V_{H_2O})/d(RH) \quad (5)$$

The LWC and surface distributions were calculated using size resolved pToF (particle time-of-flight) data of the HR-ToF-AMS. Due to the low pToF signal of NH_4 , the NH_4 surface distributions were estimated based on SO_4 pToF measurements.

2.4 Determination of OH exposure and extent of α -pinene ozonolysis

The gas-phase composition, the OH concentration and the photochemical age of a chemical reaction system may considerably differ between experiments of the same duration. Furthermore, variation of the OH concentration within a single experiment means that the photochemical age is not necessarily directly proportional to the light exposure time. Consequently, we discuss reaction time in terms of OH exposure ($\text{molecules cm}^{-3} \text{ h}$), defined as the OH concentration integrated over time. OH exposures were derived based on the decay of the OH tracer butanol-d9, detected by the PTR-MS as $M+H^+-H_2O$ at m/z 66, following the methodology introduced by Barmet et al. (2012). The OH exposure was determined by the integration of the following expression:

$$\text{OH exposure} = - \int_{t_1=0}^t \left(\frac{1}{k_{\text{OH, butanol-d9}}} \cdot \left(\frac{\Delta \ln(\text{butanol-d9})}{\Delta t} + \frac{f_{\text{dil}}}{V} \right) \right) dt \quad (6)$$

where butanol-d9 and $k_{\text{OH, butanol-d9}}$ ($= 3.4 \times 10^{-12} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$) are the butanol-d9 concentration and its reaction rate constant against OH, respectively, t is the time after lights on, V the chamber volume (we assume as a first approximation a constant chamber volume of 27 m^3), and f_{dil} the dilution flow due to HONO input (Sect. 2.2).

The percentage of α -pinene reacted with OH and O_3 was derived based on Eq. (7):

$$-\frac{d(\alpha\text{-pin})}{dt} = k_{\text{O}_3, \alpha\text{-pin}} [\text{O}_3] \times [\alpha\text{-pin}] + k_{\text{OH}, \alpha\text{-pin}} [\text{OH}] \times [\alpha\text{-pin}] + \frac{f_{\text{dil}}}{V} \quad (7)$$

where $[\alpha\text{-pin}]$, $[\text{O}_3]$ and $[\text{OH}]$ denote the concentrations of α -pinene (measured by the PTR-MS at m/z 137 and 81), O_3 and OH, respectively, and $k_{\text{OH}, \alpha\text{-pin}}$ ($= 5.3 \times 10^{-11} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$) and $k_{\text{O}_3, \alpha\text{-pin}}$ ($= 8.9 \times 10^{-17} \text{ cm}^3 \text{ molecules}^{-1} \text{ s}^{-1}$) the reaction rate constants of α -pinene with OH and O_3 , respectively. ~~The production of O_3 was faster under high NO_x compared to low NO_x , due to an efficient VOC- NO_x catalytic cycle. Therefore, the lowest percentage (65%) of α -pinene reacted with OH is presented in was achieved during experiment no. 9, under high NO_x conditions (for all experiments, the percentage of α -pinene reacted with OH ranged between 65–88%, Table 1 and discussed in Section .). We also expect that the further processing of the first generation products formed via α -pinene reaction with OH or O_3 —which don't contain C=C bonds—to predominantly proceed through reactions with OH. Accordingly, we conclude that SOA compounds detected under our conditions are mainly from OH chemistry.~~

2.5 Determination of suspended and wall-loss-corrected organic masses and yields

Suspended OA mass. The suspended organic mass concentration $C_{\text{OA}}^{\text{SUS}}$ was derived by utilizing the chemical composition measurements from the HR-ToF-AMS scaled to the total volume measured by the SMPS (Fig. S3 and Fig. S4, Supplement), using compound-specific densities ($\rho_{\text{org}} = 1.4 \text{ g cm}^{-3}$, $\rho_{\text{NH}_4\text{HSO}_4} = 1.79 \text{ g cm}^{-3}$, $\rho_{(\text{NH}_4)_2\text{SO}_4} = 1.77 \text{ g cm}^{-3}$, $\rho_{\text{H}_2\text{SO}_4} = 1.83 \text{ g cm}^{-3}$).

For some experiments (9, 10, 12 and 14-17), the organic mass concentrations determined by the HR-ToF-AMS were corrected for a sub-unity transmission efficiency at the lower edge cut-off of one of the two $\text{PM}_{2.5}$ lenses employed (Fig. S3 and Fig. S4, Supplement). Additionally, for organic mass calculation, we assumed that the measured NO_3 signals are entirely related to organonitrates (RONO_2), rather than NH_4NO_3 . This assumption mainly stems from (1) the observation of the NO_3 signal in the same particle size region as OA rather than SO_4^{2-} (see size-resolved pToF data, Sect. 3.3), while inorganic nitrate would be expected to mix within an electrolyte rich aerosol and (2) the presence of NO_3 under acidic conditions, which are thermodynamically unfavorable for the partitioning of nitric acid. This is also supported by the higher $\text{NO}^+/\text{NO}_2^+$ ratios measured in the SC compared to ratios recorded during NH_4NO_3 nebulization (on average 1–2.8 times higher, Supplement Fig. S5), typically expected from organonitrates (Farmer et al., 2010). We cannot exclude that a part of the NO_3 signal originates from NH_4NO_3 . However, even attributing all detected nitrate to NH_4NO_3 would increase the calculated LWC by 1–13 % and decrease the calculated OA mass by 2–7 % only, which would not influence our conclusions. Finally, for yield calculations we assume the accuracy of the aerosol phase measurements to be 30% (Canagaratna et al., 2007).

Wall-loss-corrected OA mass. To obtain the total C_{OA} concentration corrected for losses of particles and vapors to the chamber walls, we use Eq. (8), introduced by Hildebrandt et al. (2011), based on the mass balances of the suspended organic aerosol mass, C_{OA}^{sus} , and the mass of the organic aerosols on the walls, C_{OA}^{walls} (summed up to derive C_{OA}):

$$5 \quad \frac{d}{dt} [C_{OA}^{walls}(t)] = k_{OA}^w(t) C_{OA}^{sus}(t) + \omega(t) \cdot \left(k_{OA}^w(t) C_{OA}^{sus}(t) + \frac{d}{dt} [C_{OA}^{sus}(t)] \right) \cdot \frac{C_{OA}^{walls}(t)}{C_{OA}^{sus}(t)} \quad (8)$$

Here, k_{OA}^w represents the loss rate constant of organic particles to the walls, derived by fitting the suspended organic mass concentration 5–8 h after lights were switched on in the SC, when SOA production is expected to be negligible (α -Pinene concentration < 1 ppbv, Fig. S6 and Table S2, Supplement). We determine an average loss rate of $0.13 \mu\text{g m}^{-3} \text{h}^{-1}$, corresponding to a particle half-life of 5.3 h. The average k_{OA}^w was used in the case of
 10 insufficient statistics to perform accurate fitting. In Eq. (8), ω , ranging between 0 and 1, is a dimensionless proportionality coefficient between the mass of organic vapors that partition onto the wall-deposited particles and the mass of organic vapors that partition onto the suspended particles. Here, we neglect the condensation of organic vapors onto the wall-deposited particles, i.e. $\omega = 0$, consistent with previous studies of α -pinene SOA
 15 yields, but does not influence the comparison between the experiments. Considering the second limiting case $\omega = 1$, i.e. an equal partitioning of organic vapors between the wall-deposited and suspended particles, would increase the determined SOA yields (by up to 40 %, and by 20 % on average).

Eq. (8) does not take into consideration the loss of SOA-forming ~~vapours~~vapors onto the clean Teflon walls, ~~which may of the chamber into consideration, which depends on the wall to seed surface ratio and may greatly~~ suppress SOA yields from laboratory chambers under certain conditions. These processes may be related to the vapours' reactive uptake onto walls or to their absorptive uptake . The reactive uptake of organic vapours onto chamber walls is only significant at high RH. Loss rates for important reactive gases, including glyoxal, epoxides and peroxides have been documented at different RH . While these processes may also influence reactive SOA-forming compounds under our conditions, they occur at time-scales of hours , much longer
 20 compared to the time scales of the absorptive uptake, e.g. based on recent direct measurements of vapour losses onto Teflon walls (~10 min).

The absorption of organic compounds onto the chamber walls obeys Henry's law and depends on the compounds' accommodation coefficients and their activity at the wall/gas interface . The dependence of the compounds' absorption on RH (due to a change in accommodation coefficients or in the activity of the wall absorbed compounds) has not yet been reported to the best of our knowledge and indeed merits further investigations that are beyond the scope of the current study. Nevertheless, we ~~We~~ believe that vapour absorption onto the walls is unlikely to be ~~though that gas wall partitioning does not~~ significantly affected by RH, due to the hydrophobic nature of Teflon and its minor interaction with water under subsaturation conditions (RH < 80 %). In our case, we have maintained chamber conditions during our experiments such that vapour wall losses and their inter-experimental
 25 influence the interpretation of our results, focused on relative differences can be minimized as much as possible. This is done by (1) maintaining a relatively constant wall-to-seed surface ratio for all experiments to avoid systematic biases between experiments and (2) increasing SOA production rates, which rapidly provide a significant particle condensational sink into which condensable gases can partition. We also ~~in the yields determined under different conditions. It is also worthwhile to~~ note that vapour

~~wall losses were (1) this effect was~~ found to be minor for the α -pinene SOA system where SOA formation is dominated by quasi-equilibrium growth (Zhang et al., 2014; Nah et al., 2016), ~~(2) SOA mass formation rapidly evolved after lights were switched on and (3) we maintained a relatively constant wall to seed surface ratio for all experiments. Hence losses of organic vapors may lead to a systematic negative bias in the determined yields, but do not influence the comparison between the experiments.~~

SOA yields determination and parameterization. SOA mass yields, Y (dimensionless quantity), are calculated from Eq. (9), as the organic mass concentration formed, C_{OA} , per precursor mass consumed, $\Delta C_{\alpha\text{-pinene}}$:

$$Y = \frac{C_{OA}}{\Delta C_{\alpha\text{-pinene}}} \quad (9)$$

For comparison purposes, Y values are reported in Figure 2 and Table 1 and discussed in Sect. 3 at an OH exposure of $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$, reached during all experiments. The parameterization of smog chamber SOA yields measured is based on the absorptive partitioning theory of Pankow (Eq. (10)), which expresses the production of a set of semi-volatile surrogate products (total number N) as a function of the mass yield of these products, α_i , and their partitioning coefficients, ξ_i (a dimensionless quantity reflecting the condensed-phase mass fraction of these products). The critical parameters driving the partitioning of these products are their effective saturation concentration, C_i^* , and the total concentration of the absorptive organic phase, C_{OA} . As discussed below, we consider the absorptive organic mass as the sum of the total OA concentration and the liquid water in this phase (Sect. 3.3).

$$Y = \sum_i^N \alpha_i \xi_i = \sum_i^N \alpha_i \left(1 + \frac{C_i^*}{C_{OA}}\right)^{-1} \quad (10)$$

In Eq. (10), C_i^* (in $\mu\text{g m}^{-3}$) is a semi-empirical property (inverse of the Pankow-type partitioning coefficient, $K_{P,i}$) reflecting the saturation vapor pressure of the pure constituents ($p_{L,i}^o$) and the way they interact with the organic mixture (effectively including liquid phase activity coefficients, γ_i), as expressed in Eq. (11):

$$C_i^* = \frac{10^6 M_i \gamma_i p_{L,i}^o}{760RT} \quad (11)$$

Here, M_i denotes the compound molecular weight, R the ideal gas constant and T the temperature. Smog chamber yields from single experiments are fitted as a function of C_{OA} using the volatility basis set (VBS) (Donahue et al., 2006), which separates semi-volatile organics into logarithmically spaced bins of effective saturation concentrations C_i^* . Figure 5 shows the resulting parameterizations (lines) in comparison to the measured data (symbols) for each experiment.

To determine the yields per volatility bin (α_i), wall-loss-corrected SOA yields, Y as a function of wall-loss-corrected absorptive mass concentration C_{OA} (data presented in Figure 5) were used. We assumed a total number of 5 bins: $N = 5$ with $C_i^* = 0.01, 0.1, 1, 10$ and $100 \mu\text{g m}^{-3}$. To solve Eq. (10) for the parameters α_i , we introduced a novel approach using a Monte Carlo simulation. This approach provides best estimates of α_i values (data shown in Figure 6 and Fig. S7 in the Supplement), together with a measure for the uncertainties related to the determination of the volatility distributions from SC experiments. The calculation proceeded as follows:

1/ From the calculated yields (Y), lower (Y_1) and upper (Y_2) yield curves were determined based on the estimated measurement accuracy of C_{OA} and α -pinene mass. A possible yield domain was inscribed against C_{OA} by plotting

Y_1 and Y_2 vs. their corresponding lower and upper C_{OA} , respectively. The parameterized yield curves are shown in Figure 5.

2/ A range of possible inputs was defined for each of the parameters α_i . This range is restricted within the following interval $[0; 2*(Y_i - Y_{i-1})]$. This step was only necessary for computational reasons.

5 3/ α_i -parameters were randomly generated over the defined intervals.

4/ Deterministic computations of Y vs. C_{OA} were performed using Eq. (10) and the generated α_i inputs.

5/ Volatility distributions that fell within the domain defined in step (1) were retained, aggregated and presented as probability distribution functions for each of the five effective saturation concentrations C_i^* (probability density function, PDF, Figure 6 and Fig. S7 in the Supplement).

10 2.6 **Thermodynamic modelling Phase partitioning calculations**

General principles. Using thermodynamic modelling, we seek to understand whether observed changes in SOA bulk properties (yields and O:C ratios) with RH can be explained by ~~We introduce~~ a change in the particles' thermodynamic properties. Gas-liquid and liquid-liquid phase partitioning calculations are performed following the methods developed in Zuend et al. and Zuend and Seinfeld, using the thermodynamic group-contribution model, AIOMFAC (Aerosol Inorganic-Organic Mixtures Functional groups Activity Coefficients) to calculate activity coefficients. ~~This novel~~ approach enables predicting the phase partitioning of known organic compounds knowing their abundances in a given known mixture of organic species and electrolytes at a given RH and temperature.

~~The modelling requires the use of explicit surrogate compounds. where~~ Based on Eq. (10), the partitioning of these compounds is driven by their volatility distributions, which depends on (1) the compounds' effective saturation concentrations and (2) their relative abundances. Compounds' effective saturation concentrations, used as model inputs, were calculated based on Eq. (11) initially assuming ideal mixing ($\gamma=1$) and utilizing vapor pressures estimated using EVAPORATION. The relative abundances of these compounds in the model mixtures are based on the volatility distributions derived from experimental data ~~serve as basis for thermodynamic phase partitioning calculations of model mixtures. The methods developed by Zuend et al. (2008; 2010; 2012) and Zuend and Seinfeld (2013) (section 2.5).~~

~~We were used to calculate gas particle and liquid-liquid phase partitioning. In the model, we~~ assumed instantaneous reversible absorptive equilibrium of semi-volatile organic species into ideal and non-ideal liquid phase aerosols. ~~To calculate activity coefficients of the organic species as a function of the liquid particle phase composition, the thermodynamic group contribution model AIOMFAC (Aerosol Inorganic Organic Mixtures Functional groups Activity Coefficients) developed by Zuend et al. (2008; 2011) In the case of non-ideal solutions, positive~~ was utilized. Positive and negative deviations of mole fraction-based activity coefficients from a value of unity (ideal mixing) indicate the degree of non-ideality in a mixture. The activity coefficients take into account the compounds' affinity towards the solution (interactions with other organic species, electrolytes and water) and hence depend on the solution chemical composition. Therefore, the activity coefficients for the different organic species cannot be set a priori, but are calculated iteratively in the model (until convergence of the compounds' abundances in the different phases; see .

We considered cases with and without interactions between the electrolyte and the organic phases, which enables assessing the solubility of the organic compounds in the inorganic. ~~The overall loading of the smog chamber was used to calculate the seed aerosols. Seed~~ concentrations in $\mu\text{g}\ \mu\text{g}^{-1}\ \text{m}^{-3}$ ~~were, which was then~~ transformed in moles of seed per volume ($\text{mol}\ \text{m}^{-3}$) assuming equal shares of AHS and SA for acidic seeds and equal shares of AHS and AS for neutral seeds. Because interaction parameters of some organic functional groups with HSO_4^- are missing in AIOMFAC, we assumed the interactions of organic compounds with HSO_4^- and with SO_4^{2-} to be similar. For all computations, metastable supersaturated salt solutions were allowed.

~~Model compounds are only surrogates. The choice of these surrogates should reflect the wide range of volatility and hydrophilicity of SOA species. We have selected as surrogates α -pinene photo-oxidation products reported in literature, identified under different NO_x and aerosol seed conditions, covering the wide range of volatility relevant to SOA. Compound hydrophilicity instead depends heavily on the number of functional groups present in the molecule, which can be largely simulated by the compound O:C ratio. As increasing the compound O:C ratio also decreases its volatility, we have considered cases with and without fragmentation products with shorter carbon backbone chain, but high O:C ratios, expected to be representative of later generation photolysis and photo-oxidation products. This approach would effectively decouple compounds' hydrophilicity and volatility and allow scanning both properties independently. While we recognize that oxidation conditions, e.g. NO_x concentrations, may significantly alter the product distribution, we did not select different sets of products for the different conditions. This is because:~~

~~(1) such separation would implicitly suggest that the chemical composition of the few compounds reported at different conditions can be extrapolated to the bulk OA under our conditions;~~

~~(2) such separation would significantly limit the number of surrogates at each condition, increasing the sensitivity of the model to the compounds' selection;~~

~~(3) the model is less sensitive to compounds' chemical structure than to their elemental composition (see below and Li et al.), e.g. number of oxygen and carbon, which has been taken into account by including fragmentation products.~~

~~The relative abundance of the selected compounds is optimized in the model for each experiment at the prevailing RH such that the modelled and measured SOA yield and O:C ratio match. Then, the RH is modified in the model, and changes in the SOA yields and O:C ratios are compared to the observations. Model calculations were performed at an OH exposure of $(2.0 \pm 0.5) \times 10^7\ \text{cm}^{-3}\ \text{h}$. ~~Based on Eq. (10), SOA partitioning is driven by the compounds' volatility distributions, which depends on the compounds' effective saturation concentrations (C_i^*) and their relative abundance (α_i).~~~~

Simulated cases. The following simulations were performed:

1/ Case org: non-ideal partitioning including liquid-liquid phase separation, LLPS (activity coefficients calculated with AIOMFAC) of the organic compounds between gas phase and a purely organic aerosol phase neglecting the presence of the seed aerosol and using reported model compounds only;

2/ Case id: ideal partitioning (activity coefficients all set to unity, LLPS cannot occur) between gas phase and organic aerosol phase using reported model compounds only;

3/ Case sd: non-ideal partitioning including the seed aerosol to an internally mixed organic/AS aerosol using reported model compounds only;

4/ Case sdf: non-ideal partitioning including the seed aerosol to an internally mixed organic/AS aerosol including the formation of fragmented oxidation products (see below);

5/ Case orgf: non-ideal partitioning including LLPS to a purely organic aerosol phase neglecting the presence of the seed including the formation of fragmented oxidation products.

α_i -parameters determined through the Monte-Carlo simulations assume that the absorptive mass consists of the organic phase (in accordance with assumptions in chemical transport models). This assumption is violated for cases sd and sdf as compounds may partition into the inorganic phase. Nevertheless, the results obtained for these cases may still be examined in relative terms to inspect the effect of RH on SOA yields in the presence of an inorganic seed and the partitioning of SOA compounds between the inorganic and organic phases.

Model compounds. To simulate SOA partitioning in AIOMFAC, α -pinene photooxidation products reported in the literature (Eddingsaas et al., 2012; Jaoui and Kamens, 2001; Kleindienst et al., 2007; Valorso et al., 2011) were chosen as model compounds for cases org, id and sd, namely ValT4N10 (10th compound in Table 4 from Valorso et al. (2011)), 3-hydroxyglutaric acid, pinic acid, hopinonic acid, norpinic acid, 2-hydroxyterpenylic acid, 10-oxopinonic acid, and 4-oxopinonic acid. These compounds are listed in Table S4 in the Supplement together with their relevant physicochemical properties (MW, O:C ratio, vapor pressures and chemical structures). For cases sdf and orgf additional oxidized fragmented products (3-oxoadipic acid, glutaric acid, 5-COOH-3-OH-pentanal, and succinic acid) were included (Table S4, Supplement). While these compounds were not reported to derive from α -pinene oxidation, their structure, including carbon and oxygen numbers are very similar to the most abundant compounds detected by Chhabra et al. (2015) and Mutzel et al. (2015) using chemical ionization mass spectrometry. Volatility distributions could reliably be determined for volatility bins $C^* = 0.1$ -100 $\mu\text{g m}^{-3}$. Nevertheless, also lower volatility products with $C^* = 0.01 \mu\text{g m}^{-3}$ are formed. Therefore, this bin was loaded for all phase partitioning calculations with equal fractions of the model compounds diaterpenylic acid acetate, 3-MBTCA and ValT4N9 with mass yields of 10^{-5} each. This low mass fraction does not influence the organic yield but proved to aid the convergence of the phase partitioning calculation. To achieve mass closure in the model, pinonaldehyde (MW: 168 g mol^{-1}) was assumed to represent the more volatile products, which do not partition to the condensed phase. Because pinonaldehyde resides almost totally in the gas phase, it was not explicitly modelled.

Model compounds C_j^* calculation. Model compounds j were assigned to the volatility bins based on the calculated C_j^* values using Eq. (11) and assuming ideal mixing ($\gamma_j = 1$). Using the actual activity coefficients calculated with AIOMFAC is not possible because they are a result of the phase partitioning calculation. However, because activity coefficients proved to be generally in the range of 0.1 to 10, they did not alter the initial product assignments to the volatility bins which cover one order of magnitude in effective saturation concentration. Therefore, the initial allocation of the compounds to the volatility bins in the VBS remains valid after taking the non-ideality into account. $p_{L,j}^o$ used in the calculations were vapor pressures of pure compounds in liquid or subcooled liquid state, at 298 K estimated using EVAPORATION (estimation of vapor pressure of organics, accounting for temperature, intramolecular, and non-additivity effects, (Compernelle et al., 2011)), without using the empirical correction term for functionalized diacids.

Setting model compounds' relative abundances in the model. From the α_i -parameters generated through the Monte Carlo simulations in Sect. 2.5, 10 sets per experiment were randomly selected for the phase partitioning calculations (provided that these parameters fall within the 10th and the 90th percentiles, to avoid outliers). The 11 chosen experiments exclude experiments 15-17 with hydrophobic seeds, experiment 6 which has no counterpart experiment at high RH and experiments 5 and 14 where the COA concentration range was very limited to accurately derive α_i -parameters (Sect. 2.5). The chosen parameters are listed in Table S5 in the Supplement. As several model compounds are assigned to a volatility bin i , the yield of a compound j is expressed as its relative contribution within the volatility bin i , $\chi_{j,i}$, times the relative abundance of the bin α_i . For each experiment and each simulated case the fitted $\chi_{j,i}$ values are listed in Tables S6-S8 in the Supplement. For case org the $\chi_{j,i}$ -values were optimized to match the experimental organic yields and the measured O:C at the actual RH of the experiment. RH was then changed in the model ~~to the value of the corresponding experiment (same experimental conditions but different RH)~~ and the effects of RH on SOA yields and ~~degree on degrees~~ of oxygenation (~~O:C ratios~~) were ~~then~~ evaluated. For cases id and sd the same $\chi_{j,i}$ were used as for case org. For cases sdfr and orgfr, more fragmented compounds were added and their $\chi_{j,i}$ -values were optimized to achieve agreement between measured and modelled organic mass yields and O:C ratios. Likewise, the effects of RH on SOA yields and on O:C ratios were then evaluated by changing the RH in the model.

The following pairs of corresponding experiments were simulated: experiments 3 and 4 (66 % RH, 29 % RH), experiments 7 and 11 (67 % RH, 26 % RH), experiments 8 and 13 (60 % RH, 25 % RH), and experiments 10 and 12 (50 % RH; 26 % RH). Equilibrium calculations were performed starting from both, low and high RH experiments. For experiments 5 and 14 performed at low RH values, the volatility distribution parameters could not be determined for the most volatile bin (α_5). Therefore, equilibrium partitioning calculations were only performed starting from their corresponding experiments, namely experiment 1 (69 % RH), experiment 2 (67 % RH), and experiment 9 (56 % RH).

3 Experimental results

Figure 1 shows the suspended and wall-loss-corrected organic mass concentrations for all experiments listed in Table 1 as a function of OH exposure. SOA mass is rapidly formed and the wall-loss-corrected mass reaches a plateau at an approximate OH exposure of $2 \times 10^7 \text{ cm}^{-3} \text{ h}$. In the following, comparisons between the different experiments are carried out at an OH exposure of $2 \times 10^7 \text{ cm}^{-3} \text{ h}$. For illustrative purposes, the wall-loss-corrected aerosol yield is shown in Fig. S8, Supplement, as a function of α -pinene reacted to demonstrate its variability between different experimental conditions.

3.1 SOA yield dependence on RH, NO_x/α -pinene and aerosol seed composition

In Figure 2, we examine the relationship between the determined yields and the prevailing experimental conditions: RH, NO_x/α -pinene and seed composition. Lines connect experiments conducted under comparable conditions, but at different RH (dashed lines) and at different NO_x/α -pinene (solid lines). The visible effects of RH and NO_x conditions on the yield in Figure 2 were statistically examined using a multilinear analysis (Fig. S9 and Table S3, Supplement). Three yield parameterizations were inter-compared and only the simplest model which represented significantly better the observations was considered for discussion.

The following features can be deduced from the analysis:

1/ SOA yields increase with α -pinene concentrations, consistent with the semi-volatile nature of SOA compounds formed. We estimate that yields increase by approximately 2 percentage points% when α -pinene concentrations increase from 20 to 30 ppbv.

5 2/ A significant effect of the seed initial concentrations (or surface) on SOA yields was not observed. This may suggest (1) that SOA most likely forms its own phase and does not significantly partition into the seed aerosol (see below) and (2) that SOA condensation is not significantly affected by the vapor losses to the SC walls which can be diminished by increasing the aerosol surface. This is consistent with recent smog chamber results suggesting that for the α -pinene system SOA formation is dominated by quasi-equilibrium growth and
10 vapor losses to the walls do not depend on the seed concentrations, but rather on SOA (precursor) formation (oxidation) rates (Nah et al., 2016).

3/ SOA yields are significantly reduced under high NO_x conditions ($-3.3 \% \pm 0.6 \%$, $p < 0.001$), in agreement with literature data (Ng et al., 2007). Such a decrease indicates that SOA compounds formed under low NO_x conditions are less volatile than those formed under high NO_x conditions.

15 4/ We observed a clear influence of the RH on the yields, which increase on average by $1.5 \pm 0.6\%$ per 10% RH, for the range explored. This indicates that the particulate water content plays a central role in the partitioning of SOA compounds, either by altering the thermodynamic properties of the bulk phase (e.g. increasing the absorptive mass or decreasing the compound activity coefficients, non-reactive uptake) or by providing a reactive sink for semi-volatile species (e.g. formation of lower volatility compounds/oligomers in the
20 bulk phase, reactive uptake). Additionally, the multilinear analysis suggests that the magnitude of the RH influence on SOA depends on the seed chemical nature (yields correlate with the interaction term between RH and seed composition), with a greater influence for the acidic seed ($0.21 \pm 0.03\%$ per 1% RH, $p < 0.001$), the most hygroscopic aerosol, compared to the non-acidic seed ($0.15 \pm 0.03\%$ per 1% RH, $p < 0.001$) and the hydrophobic seed ($0.09 \pm 0.04\%$ per 1% RH, $p = 0.05$).

25 Overall, these results highlight the sensitivity of SOA yields to the prevailing oxidation conditions and the particle bulk-phase composition and water content and therefore the need for considering such conditions to obtain an accurate prediction of the SOA burden in the atmosphere. Nonetheless, results from this analysis should be regarded with some level of caution, owing to the limited size of the dataset. Despite the methodical assessment of the significance of the multilinear analysis results, we cannot unambiguously propose a
30 mechanism by which water and seed hygroscopicity/acidity enhance SOA yields nor determine whether there is interplay between these two parameters. Nevertheless, we provide evidence that both of these parameters play a significant role in the formation or condensation of SOA species. In Sect. 4, using phase partitioning computation and AIOMFAC, we shall assess to which extent and under which conditions particulate water may alter the organic species' activity coefficients and as a consequence their absorptive partitioning.

35 3.2 SOA elemental composition

The effect of the experimental conditions on SOA chemical composition was investigated using the Van Krevelen space (Figure 3a). The overall region of the experimental data is very comparable for all experimental conditions and comparable to ambient data (Ng et al., 2011a). Data for experiments 1-14, with inorganic seeds,

follow a similar slope with aging (-0.84): An increase in O:C during aerosol aging takes place under all conditions. The O:C and H:C ratios of SOA produced with hydrophobic seed aerosol (expts. 15-17) show lower values than with inorganic seeds.

In Figure 3b, data are separated according to the wall-loss-corrected organic mass concentrations C_{OA} (left panel: 2–6 $\mu\text{g m}^{-3}$; middle panel: 8–14 $\mu\text{g m}^{-3}$; right panel: 16–30 $\mu\text{g m}^{-3}$), to isolate the NO_x and RH effects on the chemical composition from the possible influence of enhanced partitioning of semi-volatile organic species to the aerosol phase due to a higher SOA loading (Pfaffenberger et al., 2013). We observe that NO_x levels have the highest influence on SOA elemental composition; namely, SOA formed at low NO_x (marked with asterisks) is characterized by a higher H:C than that formed at high NO_x . This is even more pronounced at an early reaction stage (OH exposure $< 2 \times 10^6 \text{ cm}^{-3} \text{ h}$, not shown in Figure 3), where fresh SOA products formed during all low NO_x experiments have a higher H:C and lower O:C. Such influence of NO_x levels on SOA elemental composition is consistent with our general understanding of gas-phase chemistry: Under low NO_x , substantial amounts of hydroperoxides and alcohols would result in a higher H:C ratio than e.g. carbonyls formed under high NO_x conditions (Valorso et al., 2011). Also, as hydroperoxides and alcohols are substantially less volatile than carbonyls and less fragmentation is expected at lower NO_x conditions, less oxygenated species with higher carbon number may partition to the particle phase at low NO_x , which would lead to lower O:C ratios.

NO_x levels also affect the amount of organonitrate formed. Assuming that the entire nitrate signal arises from organonitrates (i.e. a maximum organonitrate contribution), we estimate molar ratios (Fig. S10, Supplement) of NO_3 to carbon of $\sim 1:30$ for high NO_x and $\sim 1:100$ for low NO_x . Assuming oxidation product molecules with 10 carbon atoms, these ratios imply that every 3rd and every 10th molecule contains one NO_3 functional group, for high and low NO_x conditions, respectively.

Conversely, we could not observe a significant effect of RH and seed composition on SOA elemental composition and degree of oxygenation, despite their significant influence on SOA yields. This may be due to the limited range of O:C ratio spanned by the different experiments in our case (on average $\text{O:C} \in [0.56 - 0.75]$ at an OH exposure of $2 \times 10^7 \text{ cm}^{-3} \text{ h}$). Furthermore, we did not observe any significant dependence between these parameters and the ratio of organic fragments larger than m/z 150 to the total organic mass, a proxy for oligomers measured with the HR-ToF-AMS, which would explain the observed increase in yields with the aerosol liquid water content (Fig. S11, Supplement).

3.3 Analysis of aerosol size-resolved chemical composition

As the seed composition and particulate water content appear to greatly influence SOA yields, we examine in this section the interaction between these parameters and SOA, through the investigation of the aerosol size-resolved chemical composition. This information is used later to infer the behavior of the absorptive organic phase and its mixing with the inorganic seed, while modeling the aerosol dynamics in the SC is beyond the scope of this study. Figure 4 shows the aerosol size resolved chemical composition at $0 \times 10^7 \text{ cm}^{-3} \text{ h}$, $(0.5 \pm 0.2) \times 10^7 \text{ cm}^{-3} \text{ h}$, $(1.0 \pm 0.3) \times 10^7 \text{ cm}^{-3} \text{ h}$ and $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$ OH exposure for five experimental conditions (Data for additional experiments are available in the Supplement, Fig. S12). Figure S13 in the Supplement shows a 3-dimensional representation of the time dependent number and volume size distributions, measured by the SMPS.

For all experiments, the aerosol size distributions show two externally mixed aerosol populations, with a mode at lower diameters ($\sim 200 \text{ nm}$, mode 1) mostly containing SOA and another mode at higher diameters ($\sim 400 \text{ nm}$,

mode 2) mostly consisting of the seed. We note that the particle size distribution evolved consistently under different conditions, with the smallest seed particles growing with SOA condensation or coagulation (Fig. S12 and S13, Supplement). While we have detected an increase in particle number (Fig S14, Supplement), we note that intense nucleation events did not occur. For higher yields, the main SOA mass occurs in mode 1 and despite the sizeable increase of the yield with particulate water content (e.g. under acidic conditions and high RH), we did not observe a significant enhancement of SOA in mode 2. In addition, we did not note any statistically significant correlation between the initial seed volume and SOA yields; instead SOA growth seems to be driven by the favorable partitioning of semi-volatile species to smaller particles at an early stage of the experiment (Figure 4). Such behavior would imply that semi-volatile compounds do not additionally partition or react in the electrolyte rich phase on the timescale of this experiment, but rather the reactive or non-reactive uptake of these products onto the particles is enhanced with the increase of the initial particulate water content.

3.4 Prevalent oxidation reagent and its influence on SOA yields and chemical composition

Based on the mixing ratios of OH and ozone, we estimate that a greater part of α -pinene has reacted with OH (on average 0.78 ± 0.07). Moreover, it is worthwhile to mention that the further processing of the first generation products – which do not contain C=C bonds – would almost exclusively proceed through OH oxidation. Accordingly, we conclude that SOA compounds detected are mainly from OH chemistry, independent of the NO_x level and relative humidity.

We note that the fraction of α -pinene that reacted with OH under low RH (0.79 ± 0.07) and high RH (0.77 ± 0.07) are not statistically different (t-test, $p=0.71$), within our experimental variability. Therefore, we do not expect that differences in SOA yields observed between experiments at low and high RH to be due to a change in the prevalent oxidant. We recognize that water vapor may change the oxidation product distributions, via its reaction with the stabilized Criegee intermediates, produced upon the ozonolysis of α -pinene. However, we note that not only the fraction of α -pinene that reacts with O_3 is not substantial (in comparison with that reacted with OH), but also a major fraction of α -pinene Criegee intermediates undergoes unimolecular decomposition to form OH and does not react with water. Therefore, it is unlikely that a change in RH would sizably modify the distribution of the products formed via gas-phase chemistry.

By contrast, the fraction of α -pinene that reacted with OH is found to be sensitive to the NO_x concentration. The production of O_3 was faster under high NO_x (average $[\text{O}_3] = 35$ ppb) compared to low NO_x (average $[\text{O}_3] = 22$ ppb), due to an efficient VOC- NO_x catalytic cycle. Consequently, the fraction of α -pinene that reacted with OH under high NO_x (0.75 ± 0.06) is lower than that under low NO_x (0.83 ± 0.04). These small, but statistically significant differences (8 percentage points; t-test, $p=0.004$) in the contribution of ozone/OH to α -pinene oxidation are expected to explain a small part of the differences in SOA yields and chemical composition observed at low and high NO_x , with higher fraction of ozonolysis products under high NO_x conditions. Despite this, we expect the influence of NO_x on the fate of RO_2 to be the main driver behind the observed differences in SOA yields and chemical composition between low and high NO_x conditions (because ozonolysis products are only a minor fraction and differences between the two conditions are rather small).

4 Phase partitioning calculation results

The organic yields and O:C ratios of the phase partitioning calculations are presented in Figure 7 and Figure 8, respectively, and listed in Table S9 of the Supplement. Each panel in Figure 7 compares organic yields for the cases org, id, sd, sdf, and orgfr at the actual RH of the experiment (full color) and at the lower/higher RH of the corresponding experiment (light color). Table 2 gives the increase of organic yields from low to high RH of corresponding experiments as ratios [org yield (high RH) / org yield (low RH)].

The main objectives of the phase partitioning calculations are (1) estimating the impact of liquid water on SOA mixing properties (activity coefficients and liquid-liquid phase separation (LLPS)) and (2) determining the conditions or the potential model mixtures that can explain the observed yields and O:C ratios and their variation with RH. We note that model results are highly sensitive to the surrogates assumed, and the determination of SOA composition on a molecular level would considerably help confirming our results. Nevertheless, fitting both organic yields and O:C ratios significantly aid constraining the type of compounds that participate in partitioning (i.e. from a compound O:C ratio and vapor pressure its carbon number can be inferred). Accordingly, additional insights into the prevalent mechanisms by which the compounds form and evolve can be gained. For example, highly oxygenated compounds cannot be very volatile without significant fragmentation, whereas oligomerization leads to a significant decrease in the compounds' vapor pressure without necessarily increasing their O:C ratios.

4.1 Simulations with reported α -pinene photooxidation products (cases org, id, sd)

For case org, the contribution of model compounds ($\chi_{i,i}$) to the volatility bins at the actual RH could be optimized such that agreement was achieved between measured and calculated yields with deviations of less than 10 %, constituting a proof of concept of the applied approach. For most simulations, the $\chi_{i,i}$ -values optimized for case org were also valid for cases id and sd.

In general, simulations assuming cases org, id and sd failed in predicting the change in SOA yields with RH (Figure 7) and led to a significant underestimation of the SOA O:C ratios (Figure 8). The only exception is the pair of experiments 10/12 performed at low NO_x (1.9 ppbv) and high α -pinene levels (30.4 ppbv) for which the model could account for 87-99 % of the increase in SOA yield from low to high RH. This agreement was achieved although the model underestimated the O:C ratios by 24 % and predicted a smaller increase could not reproduce the high values of nor the change in the O:C values/ratios with RH than observed. The (the modelled O:C remained almost constant at 0.45-0.48 at low and high RH (difference of 6 %), while the measured O:C increased from 0.56 at low RH to 0.64 at high RH (change of 14 %). We also note that the use of the more recent parameterization by Canagaratna et al. would yield even higher O:C values, widening the gap between measured and modelled O:C ratios.

For the other experiments performed under low NO_x conditions, model simulations accounted for only 43-75 % of the observed yield increase with RH. Simulated O:C values ranged from 0.48 to 0.55 and failed to reproduce the observed increase from 0.57 to 0.62 for experiments 8/13 and from 0.6 to 0.64 for experiments 7/11. Likewise, for high NO_x conditions, model simulations for experiments 1/5, 2/5 and 9/14 could only account for 24-39 % of the observed yield increase from low to high RH and for 49-65 % for experiments 3/4.

Simulated O:C values remained almost constant at 0.49-0.54 for all high NO_x experiments and were considerably lower than the observed O:C ratios.

For case id the increased SOA yield at high RH, attributed to the additional partitioning of semi-volatile compounds (norpinic acid, 2-hydroxyterpenylic acid, 10-oxopinonic acid and 4-oxopinonic acid) is a direct consequence of the increased absorptive mass due to the higher water content. For cases org and sd, partitioning to the condensed phase is reduced compared to the case id as AIOMFAC predicts activity coefficients greater than 1 for higher volatility compounds, e.g. 10-oxopinonic acid and 4-oxopinonic acid. This effect is even enhanced for case sd when partitioning to the total condensed phase including the seed aerosol is simulated. Although LLPS is predicted for all simulations, the salting-out effect of AS which partitions to some degree also to the organic phase leads to a further decrease of the organic yield at high RH compared to the case id.

In summary our results suggest that with the reported model compounds for α -pinene photooxidation, the measured O:C and the increase of organic yields with RH cannot be simulated. Therefore, we explored whether the formation of fragmented and more oxidized products may explain the high O:C ratio observed and the high sensitivity of the yields to RH.

4.2 Simulations including fragmented products (cases sdf_r and org_{fr})

Non-fragmented products (e.g. highly oxygenated C₁₀ or dimers) would be low-volatility (LVOC) or extremely low-volatility organic compounds (ELVOC) (Zuend and Seinfeld, 2012; Donahue et al., 2006; 2012b) with effective saturation concentrations $C^* \leq 0.1 \mu\text{g m}^{-3}$, when their O:C ratio is as high as the observed O:C ratio. Therefore, they are expected to be in the particle phase independent of the prevalent RH and the inclusion in the model of additional amounts of these products would not help explaining the observed difference in the yields at different RH, leading to an overestimation of OA mass at low RH. This implies that to better capture the measured O:C range without overestimating SOA yields, fragmented products with shorter backbones need to be introduced, in accordance with Donahue et al. (2012a).

Low molecular weight compounds resulting from fragmentation were therefore added to the volatility bins with $C^* = 1-100 \mu\text{g m}^{-3}$, namely, 3-oxoadipic acid to the volatility bin with $C^* = 1 \mu\text{g m}^{-3}$, glutaric acid to the volatility bin with $C^* = 10 \mu\text{g m}^{-3}$ and 5-COOH-3-OH-pentanal and succinic acid to $C^* = 100 \mu\text{g m}^{-3}$. For case sdf_r, α_i -parameters were optimized assuming equilibrium partitioning of SOA to the whole condensed phase including the seed aerosol. For case org_{fr} absorption to the organic phase only was assumed. While the organic and the inorganic phases seem to form externally mixed particles based on the chemically resolved size distribution, examining both cases provides valuable insights into the impact of the presence of an inorganic aerosol seed in the system. Tables S7 and S8 of the Supplement provide the contribution of model compounds ($\chi_{j,i}$ values) to volatility bins C_i^* for these cases. The organic yields and O:C ratios for cases sdf_r and org_{fr} are shown in Figure 7 and Figure 8, respectively. In Table 2 the relative increase in SOA yields from low to high RH is given. Cases sdf_r and org_{fr} are similarly successful in reproducing SOA yields and O:C ratios. However, the variability in the modelled yields when using different volatility distributions for cases sdf_r and org_{fr} is larger than for cases org, id, and sd (Figure 7).

For experiments carried out at low NO_x conditions (experiments 7/11, 8/13 with NO_x < 2 ppbv), both cases explained the observed increase in the organic yields with RH and improved the agreement between modelled

and observed O:C ratios. For experiments 8/13, both cases accurately captured the increase in SOA yields from low to high RH with deviations < 2 %. Modelled O:C agreed very well with measurements for case sdfr with deviations $\leq 1\%$ and slightly worse for case orgfr with deviations of $\leq 5\%$. For experiments 7/11, cases sdfr and orgfr reproduced satisfactorily (deviations $\leq 11\%$) the measured yields for calculations with the volatility distributions determined for experiment 11, but overestimated the observed increase of O:C from low (O:C = 0.60) to high RH (O:C = 0.64). Using the volatility distributions determined for experiment 7, the low yields at low RH could not be achieved. For experiments 10/12, cases org, id, and sd were able to simulate the increase of organic yield with the oxidation products reported in literature but failed to reproduce the observed O:C ratios. Cases sdfr and orgfr with additional oxidized and fragmented products showed improved agreement of observed yields and reproduced successfully the observed O:C ratios.

Under high NO_x conditions, cases sdfr and orgfr could only simulate satisfactorily SOA yields and O:C ratios observed during experiments 3/4. For experiments 1-2/5 and 9/14, model and measurement agreement was less satisfactory. The model could not reproduce the low yields at low RH observed during experiments 5 and 14 and hence the strong reduction of the yields observed with the decrease in RH is underestimated (the simulations predict only 52-66 % of the observed change). Additionally, for these experiments the model underestimates the O:C ratios (average measured O:C = 0.66 and 0.68, at low and high RH, respectively). In the following, we examine and discuss the potential reasons that might explain model-measurements disagreements.

4.3 Simulations including organonitrates

We examined whether the discrepancy between modelled and measured yields may be ascribable to the selection of the model compounds at high NO_x , by introducing additional surrogates - specifically organonitrates - expected to be representative of the compounds formed under these conditions. Assuming that the NO_3 signal in the HR-ToF-AMS originates from organonitrates, every 3rd molecule should contain one ONO_2 functional group at high NO_x conditions. We tested whether adding organonitrates described by Valorso et al. (2011) would improve the agreement between measurements and observations for the high NO_x experiments. However, this was not the case (not shown).- Such sensitivity tests could not be performed for cases including partitioning to the seed aerosol because interaction parameters between organonitrates and sulfate are not available in AIOMFAC.

4.4 Simulations including higher volatility oxidation products

The Monte Carlo simulations enabled the determination of α_i -parameters for volatility bins $C^* = 0.1\text{-}100 \mu\text{g m}^{-3}$. α_i -Parameters in the volatility bin with $C^* = 1000 \mu\text{g m}^{-3}$ could not be reliably extracted. Nevertheless, SOA products belonging to this bin are present in the chamber and could partition to some extent to the condensed phase. We investigated whether the presence of substances with an effective saturation concentration of $1000 \mu\text{g m}^{-3}$ could reproduce the low SOA yields at low RH and its strong increase at high RH observed for the high NO_x experiments 1/5, 2/5, and 9/14. This was achieved by replacing a part of the substances in the volatility bin $C^* = 100 \mu\text{g m}^{-3}$ by pinalic acid, terpenylic acid and 3-2-oxopropanyloxopropanoic acid (Table S4, Supplement). While the addition of these compounds in the model could reproduce the low yields at low RH, the yields at high RH were strongly underestimated (not shown).- Accordingly, the introduction of organonitrates or higher volatility compounds in the model did not improve the agreement between modelled and observed yields

and O:C ratios (not shown here), suggesting that equilibrium partitioning alone cannot explain the strong SOA yield increase under high NO_x conditions.

4.5 Partitioning of individual components to the gas phase and condensed phases

The phase partitioning calculations do not only allow the simulation of the total organic yield and average O:C ratio of the condensed products, but also provide insights to the partitioning of individual compounds to the gas phase and the condensed phases. Phase partitioning of individual model compounds into the gas and condensed phases for case sdf is examined in Figure 9 for experiments 8/13 and in Figure 10 for experiments 3/4. Detailed results of the phase partitioning calculations are listed in Tables S10 and S11 in the Supplement. For experiments 8/13, the model predicts an LLPS into an organic-rich phase (op) and a predominantly electrolyte-like phase (ep). Overall, organic compounds are predominantly in the organic phase at both RH with $ep/op < 1 \cdot 10^{-5}$ at 25 % RH, and $ep/op \approx 0.003-0.008$ at 60 % RH. Compounds in volatility bins $C^* = 0.01$ and $0.1 \mu\text{g m}^{-3}$ are mainly present in the condensed phases, while compounds in volatility bin $C^* = 100 \mu\text{g m}^{-3}$ show preferred partitioning to the gas phase. The strongest increase in the condensed phase when RH is increased from 25 % to 60 % is observed for the model compounds assigned to $C^* = 100 \mu\text{g m}^{-3}$. Moderately oxygenated species (4-oxopinonic acid and 10-oxopinonic acid) in this bin show a moderate increase (about a factor of two) driven by the increase of the absorptive mass. This increased partitioning is limited by an increase in the activity coefficients of these compounds (for experiment 8 from 1.69 and 1.63 at low RH to 2.70 and 2.93 at high RH and for experiment 13 from 1.49 and 1.54 at low RH to 2.88 and 3.24 at high RH). Conversely, the 5-fold enhanced partitioning of the fragmented and more functionalized compounds (5-COOH-3-OH-pentanal and succinic acid) into the condensed phase at high RH is driven by the increase of the absorptive mass and the slight decrease of the compounds' activity coefficients (for experiment 8 from 0.84 and 0.51 at low RH to 0.68 and 0.43 at high RH and for experiment 13 from 0.90 and 0.56 at low RH to 0.69 and 0.44 at high RH).

For experiments 3/4, model results obtained from the parameterization of the yields in experiment 4 are highly sensitive to the assumed volatility distribution. As shown in Figure 10, when introducing volatility distributions characterized by high contributions of semi-volatile oxygenated compounds pertaining to the volatility bin $C^* = 100 \mu\text{g m}^{-3}$, the model predicts a liquid phase mixing. Note that when mixing is predicted, the model tends to overestimate the O:C ratio ($O:C = 0.81$ at high RH for the volatility distribution #2 shown in the upper panels of Figure 10) and the yields at both high and low RH. Conversely, when volatility distributions with less volatile and less oxygenated compounds are used (e.g. volatility distribution #93), lower yields and lower O:C ratios (0.69 for volatility distribution #93 shown in the lower panels of Figure 10) are obtained and LLPS is predicted.

For such O:C ratios LLPS has also been observed for experiments performed with model mixtures- (e.g. Song et al., 2012). For experiment 4, the volume of the electrolyte phase is larger than that of the organic-rich phase and there is considerable partitioning of organic compounds to the electrolyte phase ($ep/op = 2.26$). These observations illustrate that for compounds with high O:C ratios, small differences in the volatility distribution parameters can lead to totally different phase partitioning. ~~This highlights the suitability of the approach used, enabling the assessment of the modelled results to uncertainties in the volatility distribution determination.~~ Irrespective of these differences, none of the volatility distributions used could reproduce the measured yields at both high and low RH, likely due to the deficient representation in the model of the interactions between the acidic aerosol and the organic compounds - and of the chemical processes occurring under acidic/high NO_x conditions.

4.6 Modeling considerations and limitations

Model simulations were carried out by considering two α -pinene C8-C10 photo-oxidation products per volatility bin for cases org, id and sd. This number was sufficient to represent the species within one volatility bin, because the hydrophilicity of these compounds with consistently high carbon number correlates with their volatility. Therefore the sensitivity of model predictions to the contribution of these compounds to the volatility bins (described by $\chi_{j,i}$ values) is relatively low. This model setting, corresponding to a pseudo-one-dimensional VBS where the compounds' vapor pressure and degree of oxidation correlate, was proven to be insufficient for an accurate description of both SOA mass and degree of oxygenation. By contrast, for cases sdf and orgfr, which include more oxidized short chain products, hydrophilicity and volatility may be varied more independently, which introduces an additional degree of freedom and would correspond to a pseudo-two-dimensional VBS (2D-VBS), as described by Donahue et al. (2011). While we show that model predictions based on this setting are highly dependent on the volatility distribution parameters and model compounds assumed, in general the introduction of fragmented more oxidized compounds reproduced well the high observed O:C ratios and the increase of O:C and SOA yields with RH, specifically at low NO_x . The estimated volatility distributions and average carbon (C~6) and oxygen (O~4) numbers when considering fragmented products are in agreement with chemical speciation analysis previously reported for the same system (Chhabra et al., 2015). The analysis shows that for such semi-volatile products with O:C ratios in the order of ~0.6 an increase of RH from 23-29 % to 60-69 % induces a mass increase by up to a factor of three, driven by the higher particle water content and the lower activity coefficients of the more fragmented products at high RH.

The measured O:C ratio is a key parameter for constraining the model. We do not expect the accuracy of the O:C ratios determined by the HR-ToF-AMS to be less than 20-30% (Aiken et al., 2007; Canagaratna et al., 2015; Pieber et al.). Here, we have used the high resolution parameterization proposed by Aiken et al. (2007), while the use of the more recent parameterization by Canagaratna et al. (2015) ~~is now available. The utilization of the latter~~ would result in even higher O:C ratios (by 18 %). Higher O:C ratios 20%, ~~which~~ would require increasing even further the contribution or the degree of oxidation of the fragmented compounds and would imply that the model predicts anwould predict even a-higher sensitivity of the yields to RH. Therefore, the O:C values used here yield more conservative estimates of the contributions (or the degree of oxidation) of fragmented products in the model and of the sensitivity of the yields and O:C ratios to the RH.

Under high NO_x (and low pH), the model could not reproduce the factor of six increase in yields at high RH, using a variety of chemically dissimilar surrogate compounds. This indicates that the increased absorptive uptake of these compounds due to the increase in SOA mass and the decrease in the compounds' activity coefficients cannot explain the observed enhancements alone. Under these conditions, additional processes may likely play a role in the enhancement of OA with RH. NO_x concentrations have a dramatic influence on SOA chemical composition and volatility. Therefore, the discrepancy between model and measured yields at high NO_x conditions may be explained by either an inadequate representation of SOA surrogates in the model and their interaction parameters with the seed or by an enhanced reactive uptake of SOA species formed under high NO_x conditions, such as carbonyls that are reported to instigate the formation of lower volatility compounds in the particle phase (Shiraiwa et al., 2013 and references therein). Given the dearth of additional chemically resolved measurements of the particle phase species formed under different conditions, the mechanism by which RH enhances the uptake of SOA species under high NO_x (low pH) remains currently undetermined.

Cases org, id, and orgfr assumed gas-particle partitioning into the organic aerosol only. In this case, the role of the seed aerosol is restricted to providing a substrate for nucleation of organic vapors. This is the case for effloresced seed particles. For liquid seed aerosols this is equivalent to a complete organic/electrolyte phase separation with no partitioning of inorganic ions to the organic phase. For cases sd and sdf equilibrium partitioning between the gas phase and the entire condensed phase including the seed aerosol is assumed leading in most cases to LLPS. The current model does not yet contain interaction parameters of bisulfate with all involved organic functional groups. Therefore, ammonium bisulfate was treated as ammonium sulfate in the model, and we are not capable of distinguishing whether the enhanced partitioning of semi-volatile vapors in the acidic medium is attributable to additional reactions in the bulk phase (catalysed at lower pH) or to an enhanced solubility of SOA species.

Considering the ~~size-resolved particle chemical composition HR-ToF-AMS data in~~ Figure 4 discussed in Sect. 3.3 (~~), LLPS is likely not realized within single particles but manifests in the aerosol population splits up into a predominantly organic mode at ~200 nm and a predominantly inorganic mode at ~400 nm. presence of two externally mixed particle populations.~~ The formation of these two populations may occur by the homogeneous or heterogeneous nucleation of highly oxidized non-volatile products. Homogeneous nucleation implies new particle formation (~~which would only be moderate – see Figures S13 and S14 – due to the high condensation sink in the chamber~~), while heterogeneous nucleation proceeds via condensational growth (~~which would occur on smaller particles with higher surface~~). Both processes are expected to create small organic rich particles, providing an organic absorptive phase into which additional semi-volatile compounds may preferentially partition. When the organic and electrolyte phases are present in different particles the two phases communicate via gas phase diffusion, and equilibration ~~time-scales depend~~ ~~occurs depending~~ on the ~~components'~~ ~~of the components~~ volatility. For compounds with $CC_j^* = 0.1-100 \mu\text{g}-\mu\text{g m}^{-3}$ equilibration occurs within time-scales of minutes to tens of minutes, assuming no bulk phase diffusion limitations (Marcolli et al., 2004). ~~In~~ ~~It is expected that in~~ the larger particle ~~electrolyte-rich mode, mode a liquid-liquid phase separation will not establish because~~ the inorganic ions would exert a salting-out effect driving the organic compounds to partition to the gas-phase or into the smaller organic-rich particles. ~~This would prevent the organic compounds from partitioning in significant amounts into the seed aerosol from the beginning and, which would~~ deplete even further these larger particles from the organic material. Under such a scenario an externally-mixed phase-separated aerosol ~~may~~ ~~might~~ evolve in the smog chamber. Overall, the size-resolved chemical composition information confirms the modelling results providing compelling evidence for organic-electrolyte LLPS.

The scenario outlined above is based on equilibrium partitioning and does not invoke diffusion limitations within the condensed phase. Recent evidence may challenge this assumption, suggesting that SOA may adopt a highly-viscous state (e.g. Virtanen et al., 2010; Koop et al., 2011), where bulk diffusion and evaporation are kinetically limited. However, while such behavior occurs under certain conditions, e.g. low temperature and low relative humidity, we expect that this is not the case for the aerosol investigated in this study. Saleh et al. (2013) showed that SOA from α -pinene ozonolysis reaches equilibrium with the gas phase within tens of minutes at low mass loadings ($2-12 \mu\text{g m}^{-3}$) upon a step-change in temperature. Robinson et al. (2013) determined by aerosol mixing experiments that the diffusion coefficient in α -pinene-derived SOA is high enough for mixing on a time scale of minutes. Fast mixing is further supported by measurements of ambient OA (Yatavelli et al., 2014) showing that biogenic SOA reaches equilibrium within atmospheric time scales, under similar conditions as those in our

chamber. Therefore, we expect that, if thermodynamically favorable, liquid-liquid mixing would have occurred under the time-scales of our experiments and a unimodal particle population would have emerged. However, consistent with model predictions, this is not the case.

5 Summary and conclusions

5 We conducted a series of smog chamber experiments to investigate the impact of NO_x/VOC ratios, RH and the seed aerosol composition on the yield and the degree of oxygenation of SOA produced from α -pinene photooxidation. We developed a novel approach based on Monte Carlo simulations to determine SOA volatility distributions from the measured yields using the volatility basis set framework. Measured yields spanned an order of magnitude (2-20 %), depending on the prevailing oxidation conditions, the particle bulk-phase
10 composition and the water content. This underlines the need of considering these conditions for an accurate prediction of the SOA burden in the atmosphere. The yields increased dramatically with RH, aerosol acidity and at low NO_x. The aerosol bulk chemical composition measured by the HR-ToF-AMS appears to be mostly dependent on NO_x/VOC ratios.

We investigated whether equilibrium partitioning between the gas and the condensed phase(s) can reproduce the
15 measured SOA yield and O:C ratio at low and high RH. For this, the thermodynamic group-contribution model AIOMFAC was used to examine the partitioning of a range of selected surrogate compounds with different volatilities and O:C ratios as a function of their activity coefficients and the particle water content; properties that are altered with the variation of RH. In practice, two to four surrogate compounds were assigned to each volatility bin based on the volatility distributions derived from experimental data and the compounds' effective
20 saturation concentration, which depends on the pure component saturation vapor pressures. The RH was then varied in the model and changes in SOA mass and O:C ratios were monitored. There are large discrepancies between vapor pressures determined for semi-volatile and low volatility compounds depending on the measurement techniques (Huisman et al., 2013; Bilde et al., 2015), which introduce inaccuracies in saturation vapor pressure estimations. However, these uncertainties do not affect the approach applied in this study because
25 an incorrect assignment of a model compound to a volatility bin does not change the experimentally derived volatility bin distribution. The share of the individual model compounds to the volatility bins was used as fitting parameter to achieve agreement between measurements and calculations. In addition, as a set of volatility distribution functions was obtained for each experiment through the Monte Carlo simulations, the approach is proven very effective in assessing the sensitivity of equilibrium partitioning calculations on the volatility
30 distribution input parameters.

Our Modelling results show that only with the inclusion of fragmentation products the model could in order to simultaneously explain fit the SOA concentrations masses and O:C ratios, compounds arising from fragmentation need to be considered. Under these conditions, the model predicts that an increase in RH from ~25% to ~60% may lead to a three-fold enhancement in SOA mass, due to the increase in the absorptive mass and the slight
35 decrease in the compounds' activity coefficients. While the magnitude of this increase is consistent with the experimental observation at low NO_x, equilibrium partitioning alone could not explain the strong increase in the yields with RH observed for the high NO_x experiments (factor ~6). This suggests that other processes including the reactive uptake of semi-volatile species into the liquid phase may occur and be enhanced at higher RH, especially for compounds formed under high NO_x conditions such as carbonyls. Future studies should investigate

the dependence of SOA compounds on RH, NO_x and aerosol acidity at a molecular level. Such measurements shall provide additional insights into the chemical nature of the compounds that additionally partition or react at higher RH and also the mechanisms via which such processes occur.

5 For most of the cases studied, the model predicts liquid-liquid phase separation into an organic and an electrolyte phase. Considering the size-resolved particle chemical composition, this phase separation is likely not realized within single particles but the aerosol population splits up into a predominantly organic mode at ~200 nm and a predominantly inorganic mode at ~400 nm. Based on our results, such~~Such~~ liquid-liquid phase separation is expected to occur under most ambient conditions, i.e. similar levels of OA and sulfate, O:C < 0.8, and RH < 70–80 %.

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6 Tables and Figures

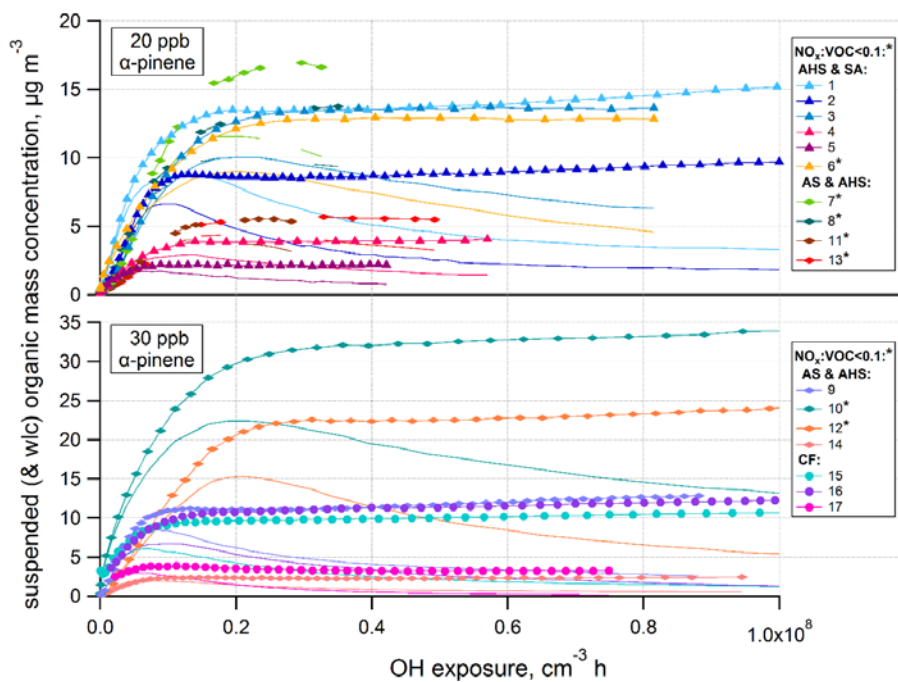
Table 1. Overview of experimental conditions. Seed types {AHS = ammonium hydrogen sulfate (NH_4HSO_4); SA = sulfuric acid (H_2SO_4); AS = ammonium sulfate ($(\text{NH}_4)_2\text{SO}_4$); CF = fluorinated carbon (see Sect. 2.2)} and their assumed phase states: (L) = liquid; (-) = liquid and/or solid. Initial seed mass concentrations; relative humidity (RH); measured mean NO_x concentrations during low NO_x experiments (marked with an asterisk*) and measured initial NO_x concentrations during high NO_x experiments; measured initial α -pinene concentrations (which reacted before an OH exposure of $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$) and their estimated fractions reacted with OH radicals in %, rest reacted with O_3 . NO_x/α -pinene ratios; wall-loss corrected organic mass concentrations (C_{OA}) and corresponding yields Y averaged over the OH exposure of $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$. Standard deviations (1sd) given in brackets are measurement variability. Horizontal lines separate experiments with different i) seed composition, ii) RH, iii) NO_x/α -pinene. Blank experiments (B1, B2 and B3) are listed at the very bottom.

No	seed type (phase)	seed initial $\mu\text{g m}^{-3}$	RH %	NO_x initial or mean(*) ppbv	α -pin initial = reacted ppbv	α -pin decay by OH %	NO_x/α -pin	C_{OA} (wlc) at OH exposure: $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$ $\mu\text{g m}^{-3}$	Yield, Y (wlc)
1	AHS+SA (L)	8.0(0.5)	69(2)	44.4(0.8)	20.7	80.0	2.1	13.4(0.2)	0.115
2	AHS+SA (L)	12.3(0.5)	67(2)	70.4(1.3)	18.7	80.3	3.8	8.6(0.1)	0.081
3	AHS+SA (L)	4.9(0.3)	66(2)	19.6(0.7)	16.1	81.6	1.2	12.6(0.6)	0.138
4	AHS+SA (L)	4.7(0.2)	29(1)	23.6(0.6)	19.9	81.3	1.2	3.9(0.0)	0.035
5	AHS+SA (L)	8.0(0.3)	28(1)	52.1(0.6)	20.3	74.6	2.6	2.1(0.1)	0.018
6	AHS+SA (L)	5.2(0.3)	27(1)	1.3(0.4)*	18.3	87.8	0.071	12.0(0.5)	0.116
7	AS+AHS (L)	8.8(0.4)	67(1)	0.7(0.2)*	20	81.9	0.037	16.2(0.4)	0.143
8	AS+AHS (L)	4.3(0.6)	60(1)	1.0(0.2)*	18.7	86.5	0.052	12.3(0.4)	0.116
9	AS+AHS (L)	5.5(0.2)	56(2)	65.8(0.8)	30.9	65.0	2.1	11.1(0.1)	0.064
10	AS+AHS (L)	4.4(0.2)	50(1)	1.3(0.2)*	30.6	79.2	0.041	29.6(1.1)	0.171
11	AS+AHS (-)	4.1(0.2)	26(1)	0.7(0.3)*	18.9	81.3	0.039	5.5(0.2)	0.051
12	AS+AHS (-)	3.6(0.1)	26(1)	1.9(0.4)*	30.5	78.4	0.062	20.3(1.3)	0.118
13	AS+AHS (-)	8.2(0.3)	25(1)	1.1(0.3)*	19.6	87.4	0.055	5.3(0.1)	0.048
14	AS+AHS (-)	3.2(0.2)	23(1)	75.1(0.7)	27.8	69.4	2.7	2.4(0.1)	0.015
15	CF (L)	7.1(0.3)	58(1)	56.2(0.7)	31.7	67.9	1.8	9.8(0.1)	0.054
16	CF (L)	10.0(0.5)	58(2)	58.3(0.5)	31.3	73.9	1.9	10.7(0.1)	0.061
17	CF (L)	6.7(0.1)	26(1)	53.3(0.6)	30.5	69.7	1.7	4.7(0.1)	0.021
B1	CF (L)	0.3(0.1)	58(2)	53.0(0.6)	-	-	-	< 0.1	-
B2	AS+AHS (L)	4.5(0.7)	68(2)	0.9(0.5)*	-	-	-	2.8(1.0)	-
B3	AHS+SA (L)	3.8(0.2)	75(3)	1.4(0.3)*	-	-	-	0.5(0.1)	-

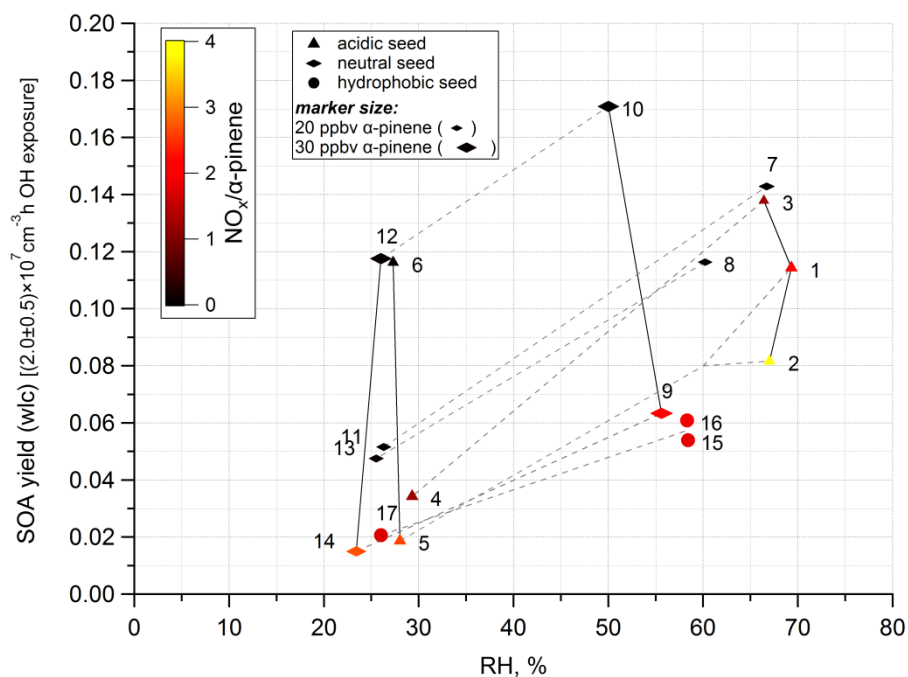
Table 2. Increase of organic yield from low to high RH for the different experiments as org yield (high RH) /org yield (low RH). The second column lists the ratio of measured organic yields enhancements with RH. In the five last columns the ratios of organic yields calculated at high and low RH for the different experiments are given (average value of the ten volatility distributions with lowest and highest values in brackets in the second row).

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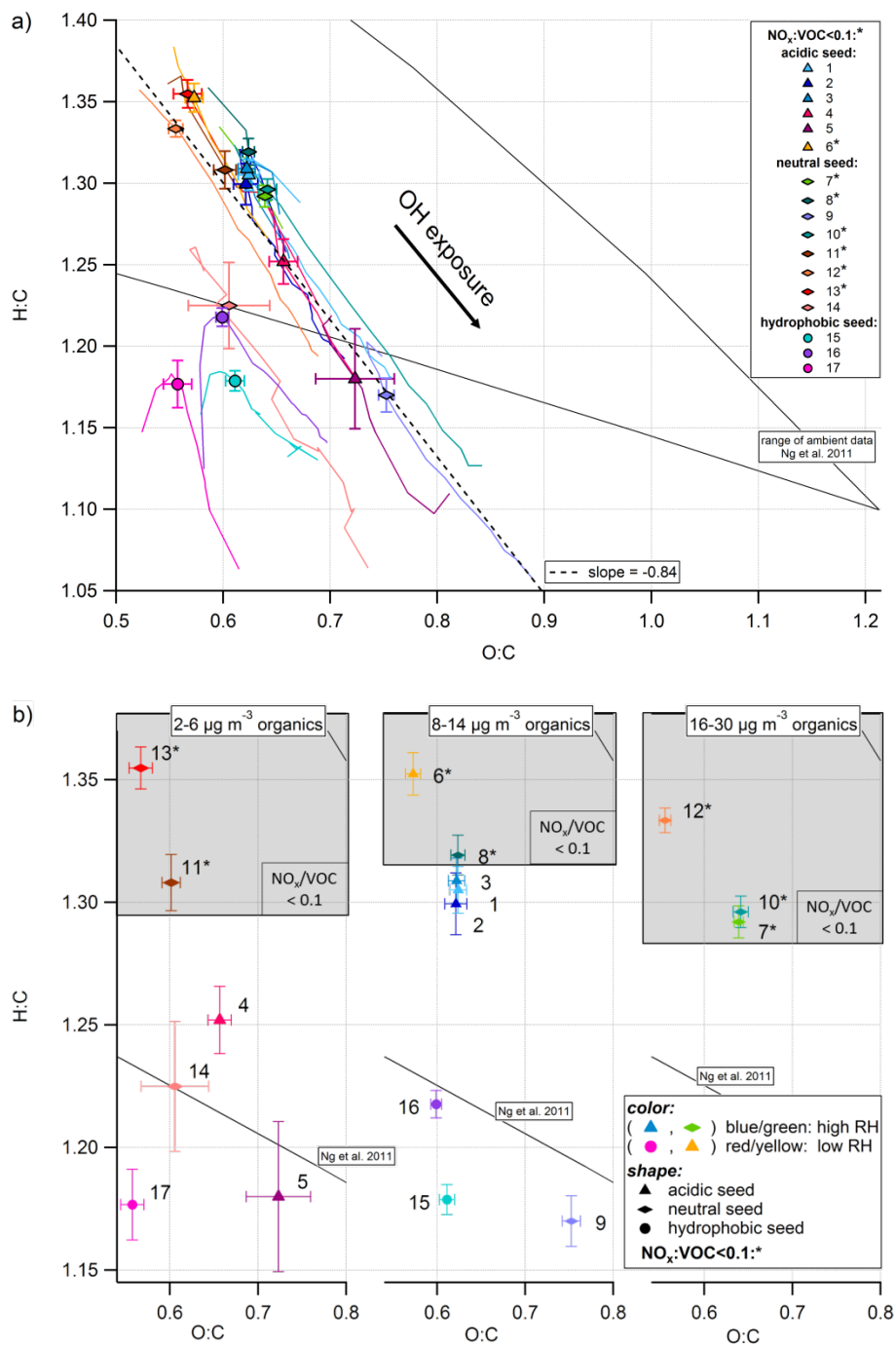
Experiments		Calculations					
Expt. No.	yield ratio	Exp	Case org	Case id	Case sd	Case sdfr	Case orgfr
1 / 5	6.38	1	1.62 (1.48-1.86)	2.02 (1.67-2.57)	1.54 (1.37-1.77)	4.23 (2.50-6.70)	4.34 (2.89-6.62)
2 / 5	4.10	2	1.42 (1.23-1.47)	1.61 (1.31-1.73)	1.31 (1.18-1.35)	2.74 (1.33-6.24)	2.24 (1.35-2.92)
3 / 4	3.23	3	1.64 (1.44-1.86)	2.07 (1.88-2.38)	1.57 (1.48-1.73)	3.67 (2.74-5.29)	3.13 (2.53-4.14)
3 / 4	3.23	4	1.76 (1.58-1.90)	2.09 (1.78-2.41)	1.58 (1.44-1.71)	2.93 (1.91-4.88)	3.12 (2.13-4.09)
7 / 11	2.95	7	1.35 (1.25-1.47)	1.52 (1.35-1.80)	1.28 (1.20-1.39)	2.16 (1.48-3.98)	2.38 (1.64-3.69)
8 / 13	2.32	8	1.43 (1.30-1.58)	1.61 (1.43-1.87)	1.36 (1.27-1.50)	2.66 (1.60-4.79)	2.44 (1.73-3.57)
9 / 14	4.63	9	1.43 (1.32-1.59)	1.56 (1.38-1.82)	1.33 (1.23-1.48)	2.45 (1.61-4.70)	2.69 (1.75-4.24)
10 / 12	1.46	10	1.29 (1.13-1.52)	1.38 (1.22-1.67)	1.28 (1.17-1.49)	1.44 (1.18-1.66)	1.55 (1.26-2.17)
7 / 11	2.95	11	1.80 (1.52-2.21)	2.22 (1.69-3.09)	1.67 (1.40-2.05)	2.74 (1.75-4.03)	3.38 (1.87-5.61)
10 / 12	1.46	12	1.34 (1.21-1.48)	1.48 (1.27-1.71)	1.33 (1.19-1.49)	1.34 (1.19-1.65)	1.70 (1.33-2.06)
8 / 13	2.32	13	1.54 (1.39-1.68)	1.74 (1.52-1.91)	1.41 (1.30-1.50)	2.73 (1.29-6.06)	2.54 (1.39-4.90)



5 **Figure 1.** 20-min-averaged wall-loss-corrected (symbols & lines) and suspended (lines) organic mass concentrations as a function of OH exposure. Data is separated according to similar initial α -pinene concentration (20 ppbv - top panel, 30 ppbv - bottom panel). Experiment numbers are given in the legend and classified by seed composition; asterisks indicate low NO_x experiments.



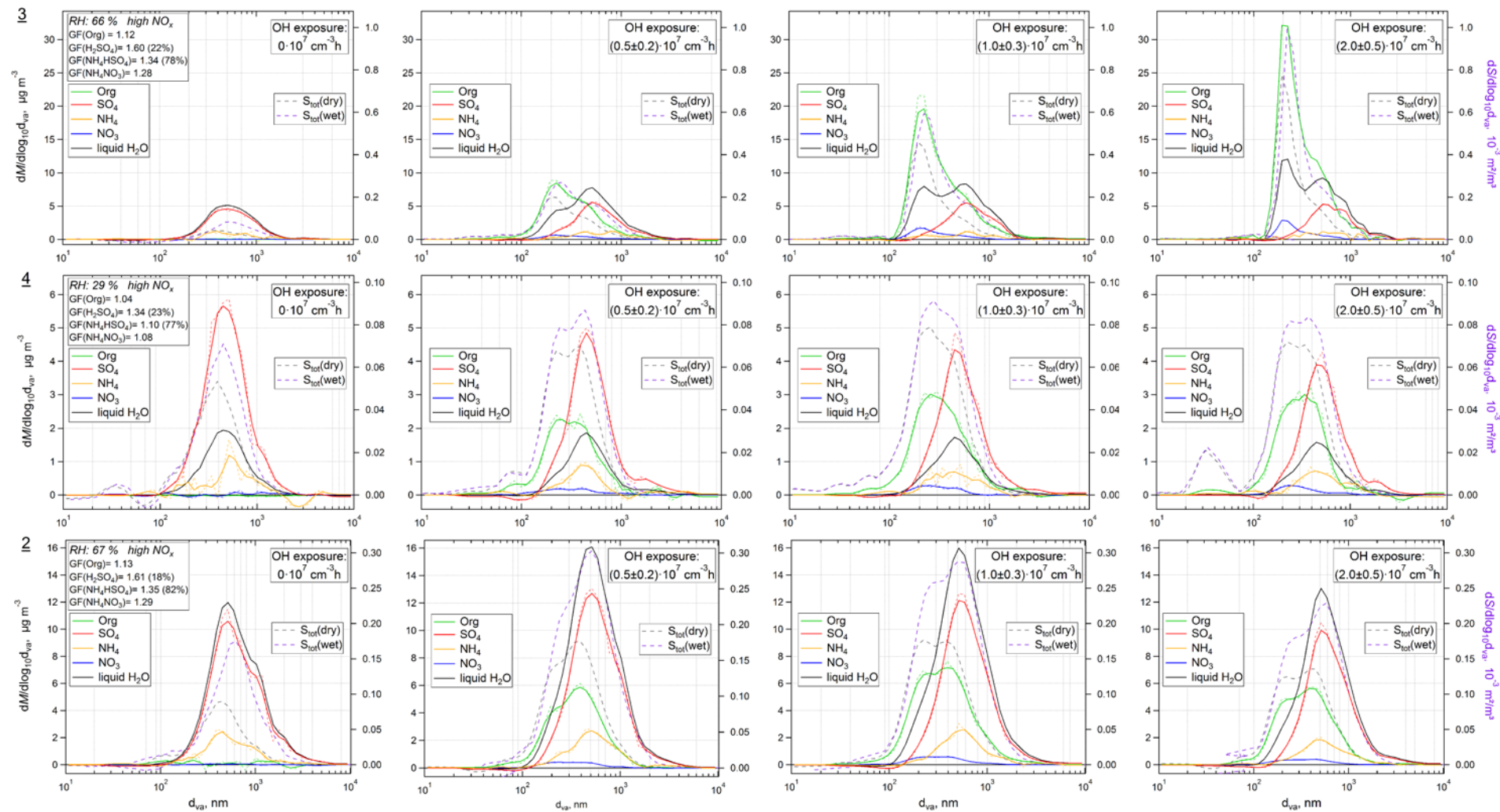
5 **Figure 2.** Average wall-loss-corrected yields Y at $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h OH exposure}$ as a function of RH. Symbol sizes represent α -pinene reacted, symbol colors represent NO_x/α -pinene and symbol shapes represent seed composition according to Table 1. Experiments with similar NO_x/α -pinene, seed composition and α -pinene reacted are connected with dashed lines showing the increase in yield for increased RH. Experiment 6 has no counterpart experiment at high RH. Experiments with similar RH, α -pinene reacted and seed composition are connected with solid lines showing the increased yield with decreasing NO_x/α -pinene ratio.



5 **Figure 3. Van Krevelen diagrams: Mean (and 1 standard deviation measurement variability) H:C versus O:C at OH exposure $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$ (symbols). Symbol colors indicate RH, symbol shapes the seed composition and the asterisk low NO_x experiments. (a) The dashed line represents the (least orthogonal distance) fitted slope of experiments (1-14) with inorganic seed: -0.84. The triangular shaped solid lines) represents the range of ambient SOA (Ng et al., 2011a). 1 h-averages of H:C versus O:C (for OH exposure $> 2 \times 10^6 \text{ cm}^{-3} \text{ h}$) are given by the colored lines. Experiments 15-17 show the influence of organic seed compounds (data shown from suspended organic mass $> 0.3 \mu\text{g m}^{-3}$). (b) The data was split in three groups according to their wall-loss-corrected organic mass concentrations to exclude concentration effects (left panel: $2\text{-}6 \mu\text{g m}^{-3}$; middle panel: $8\text{-}14 \mu\text{g m}^{-3}$; right panel: $16\text{-}30 \mu\text{g m}^{-3}$). The grey shaded areas include all low NO_x experiments.**

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a)



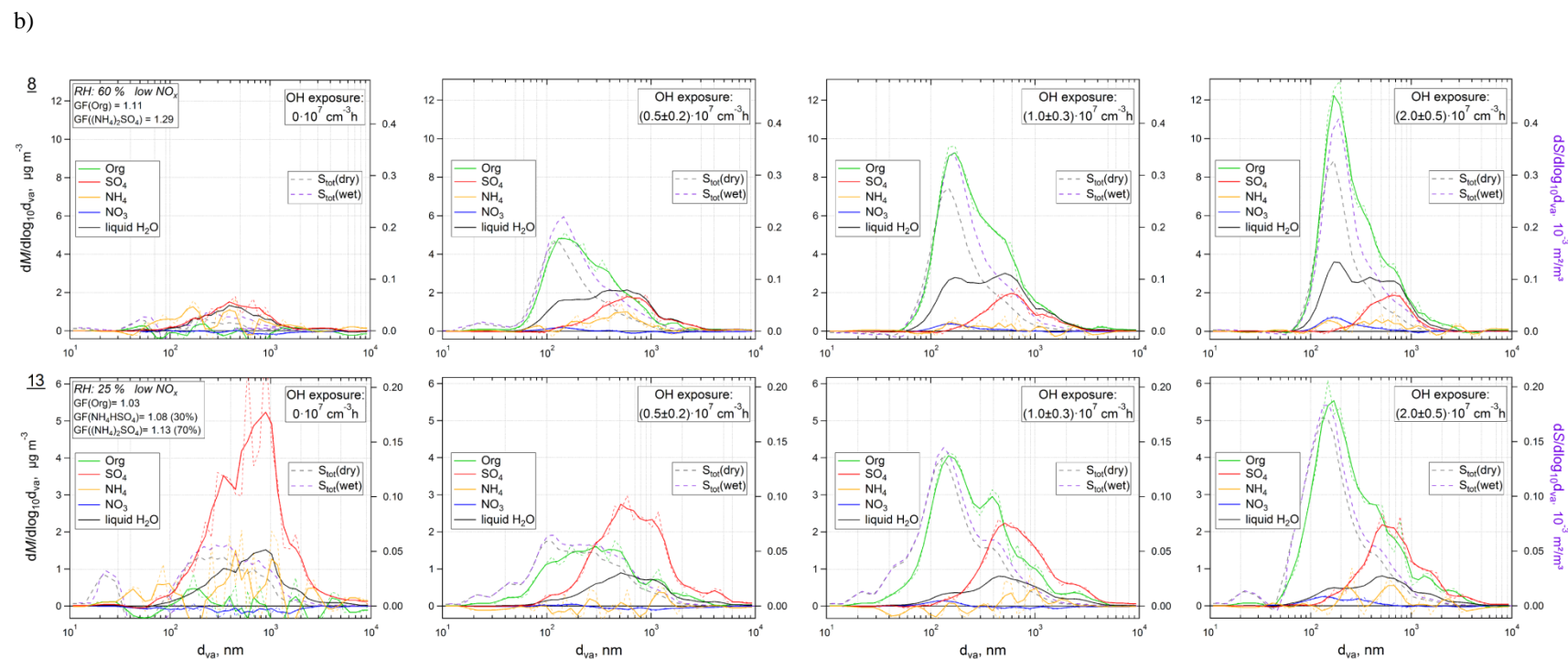
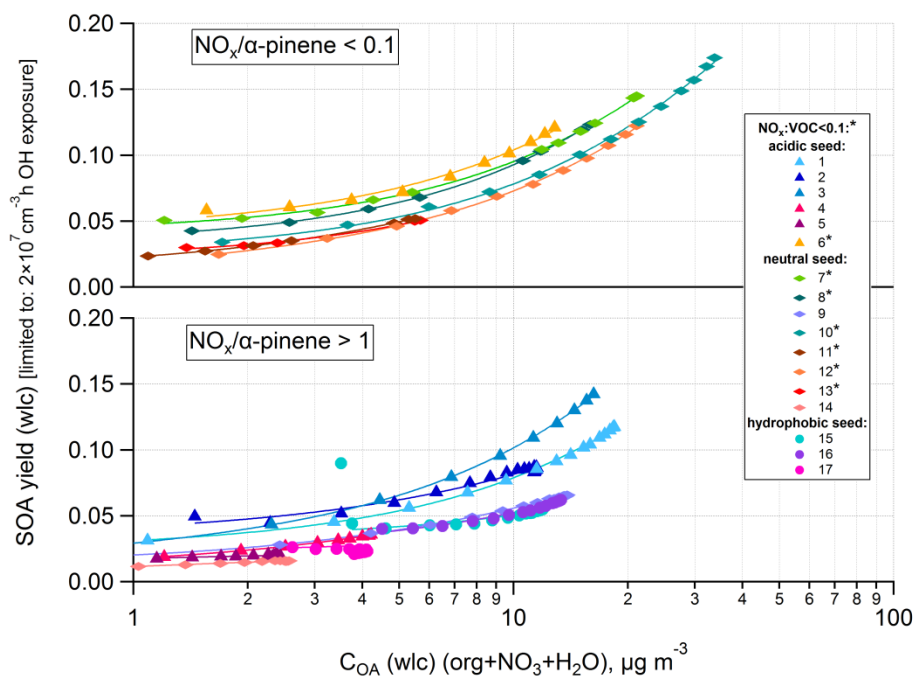
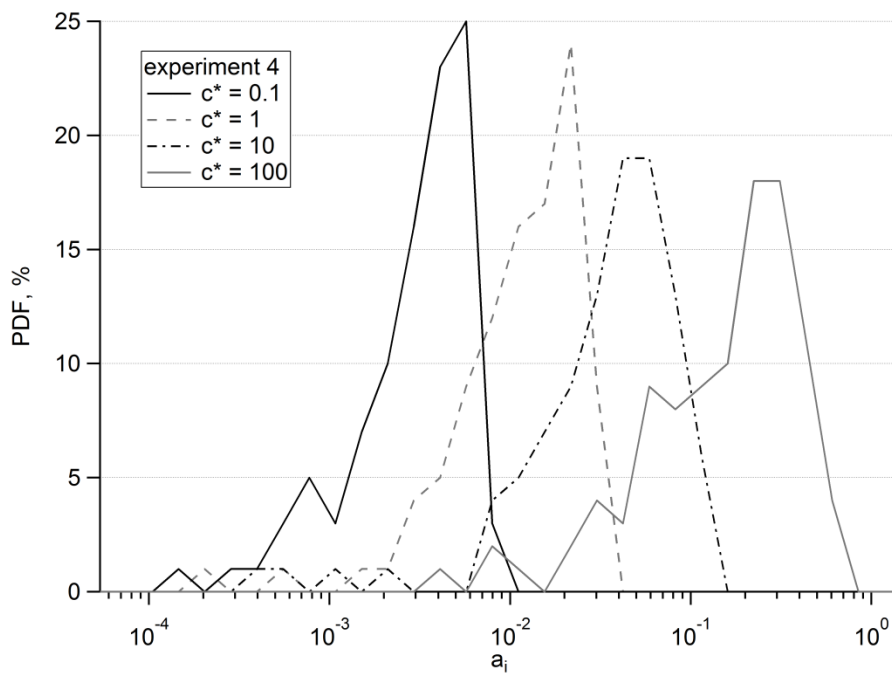
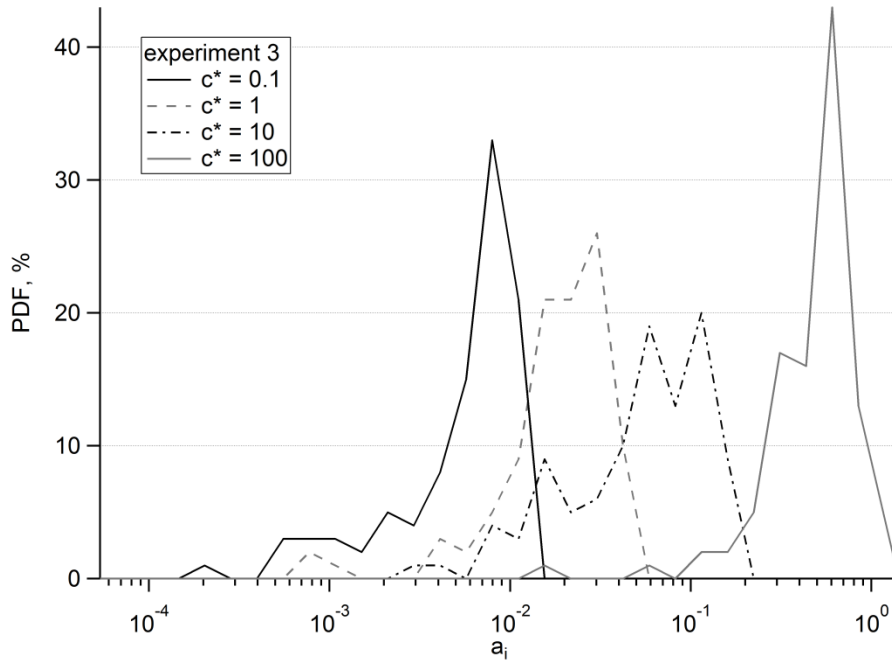


Figure 4. Evolution of size distributions from the HR-ToF-AMS. Measured organic, SO_4 , NH_4 , NO_3 mass distributions for OH exposures of $0\times$, $(0.5 \pm 0.2)\times$, $(1.0 \pm 0.3)\times$ and $(2.0 \pm 0.5) \times 10^7 \text{ cm}^{-3} \text{ h}$. Black lines represent estimated liquid water content (method: Sect. 2.3; RH and individual GFs given in the legend, percentage in brackets: fractions of SO_4). The calculated dry and wet surface distributions are shown as dashed lines on the right axes. (a) High NO_x with more acidic seed [H_2SO_4 & NH_4HSO_4]: Experiments 3 and 4 with NO_x/α -pinene = 1.2 and experiment 2 with NO_x/α -pinene = 3.8. (b) Low NO_x with less acidic seed [$(\text{NH}_4)_2\text{SO}_4$ & NH_4HSO_4]: Experiments 8 and 13 with NO_x/α -pinene ≈ 0.05 . Additional figures in the Supplement (Fig. S12).

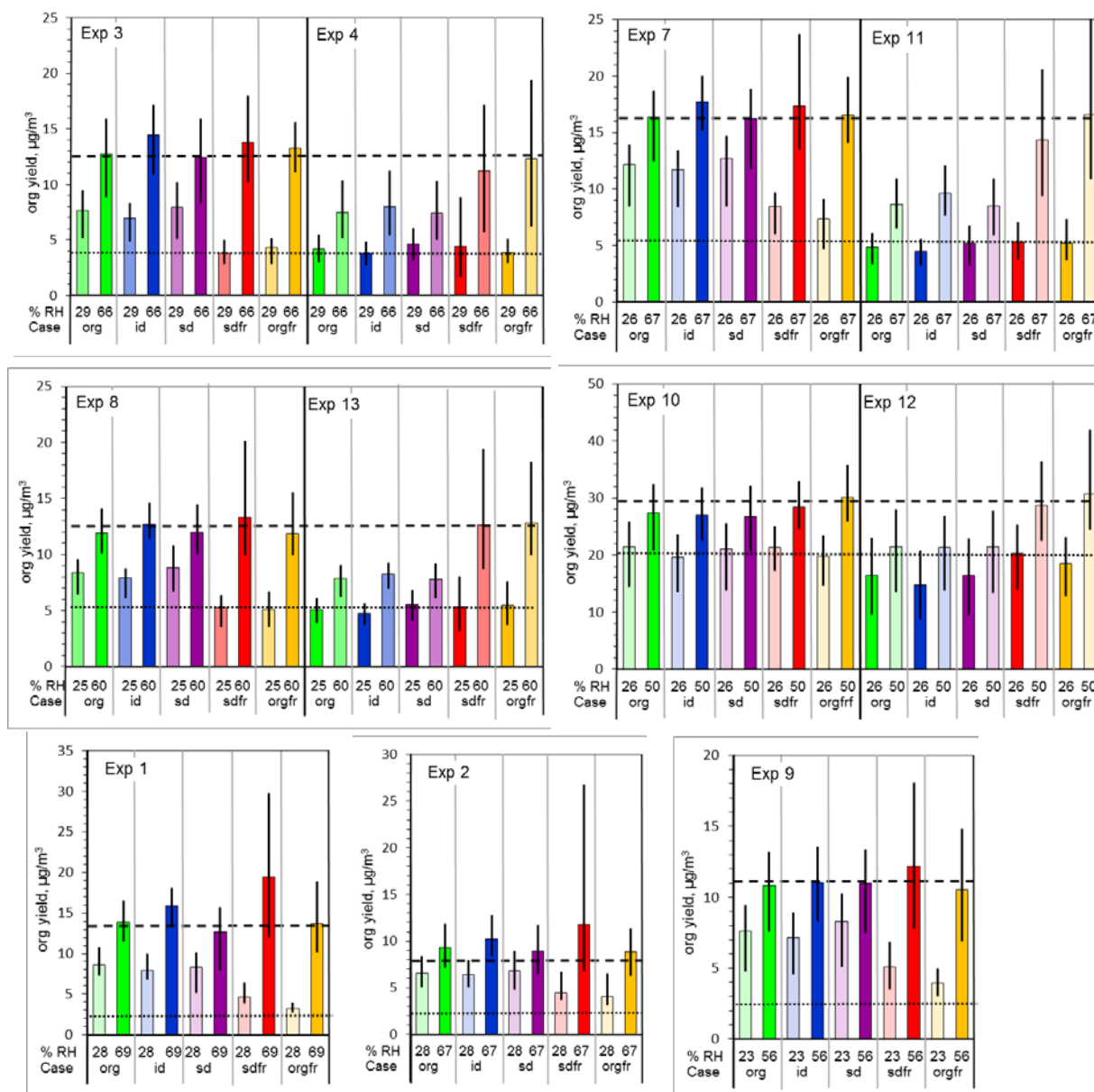


5 **Figure 5.** Measured (symbols) and parameterized (lines) wall-loss-corrected SOA yield, Y , as a function of wall-loss-corrected absorptive mass concentration (C_{OA} : organics + NO_3 + H_2O) for low NO_x experiments (upper panel) and high NO_x experiments (lower panel). Data was limited to an OH exposure of $2 \times 10^7 \text{ cm}^{-3} \text{ h}$ and averaged over 20 minutes. Symbol colors indicate the RH, symbol shapes the seed composition and the asterisks the NO_x/VOC ratio.



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Figure 6. Probability density functions (PDF) of α_i 's for volatility bins ($C_i^* = 0.01, 0.1, 1, 10$ and $100 \mu\text{g m}^{-3}$) for one high RH (exp. 3, upper panel) experiment and one low RH (exp. 4, lower panel).



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Figure 7. Calculated average organic yields for experiments at the measured RH (columns in full colors) and at the RH of the corresponding experiments (light colors, see Table 2 for reference) based on ten randomly chosen volatility distributions for the cases org, id, sd, sdfr, and orgfr (identified by different colors). The vertical black lines on the columns indicate the range of values obtained for the calculations with the individual volatility distributions. The horizontal dotted line marks the measured value for the experiment performed at low RH, the dashed line the value for the experiment performed at high RH.

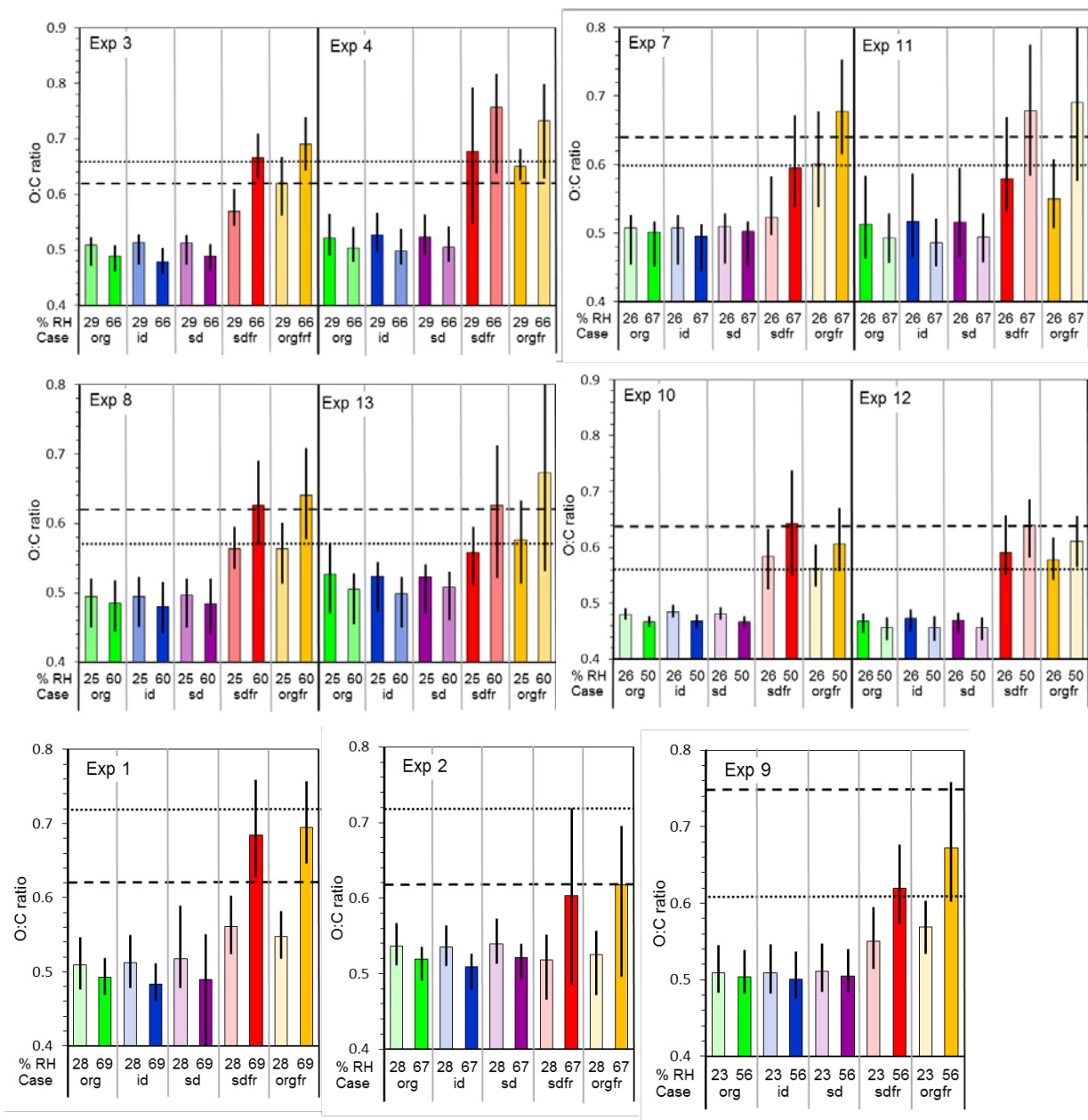
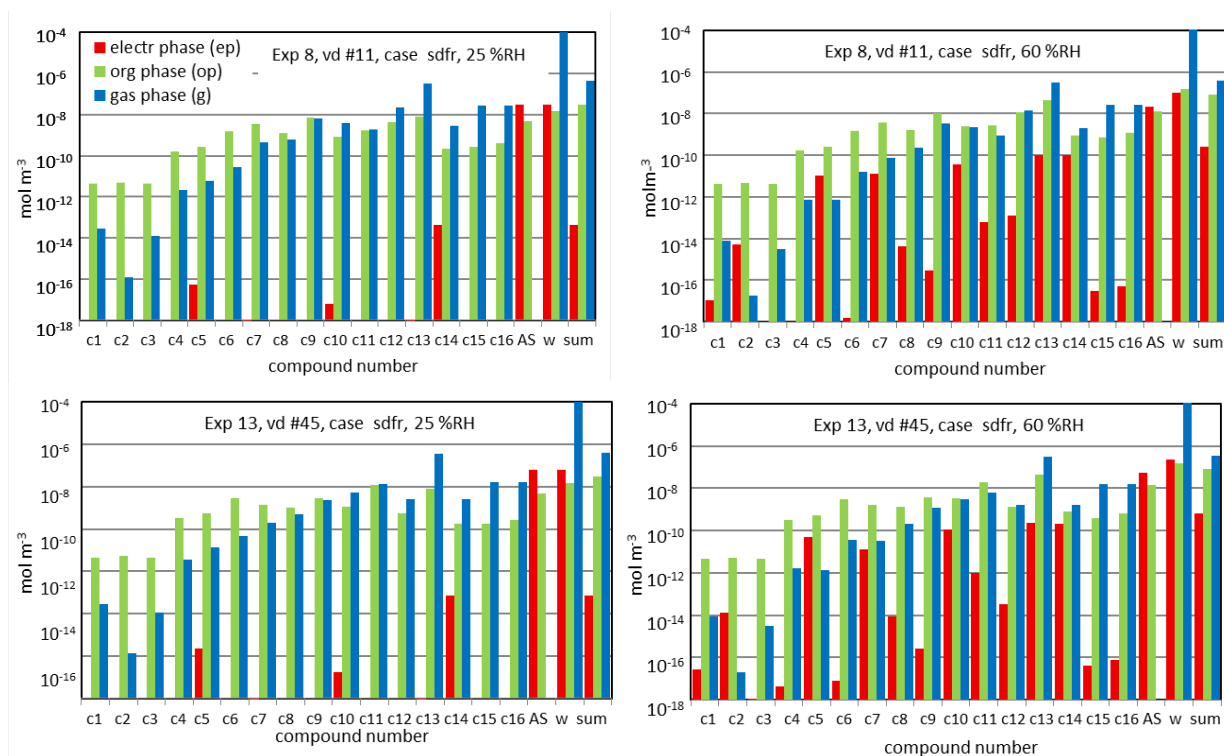
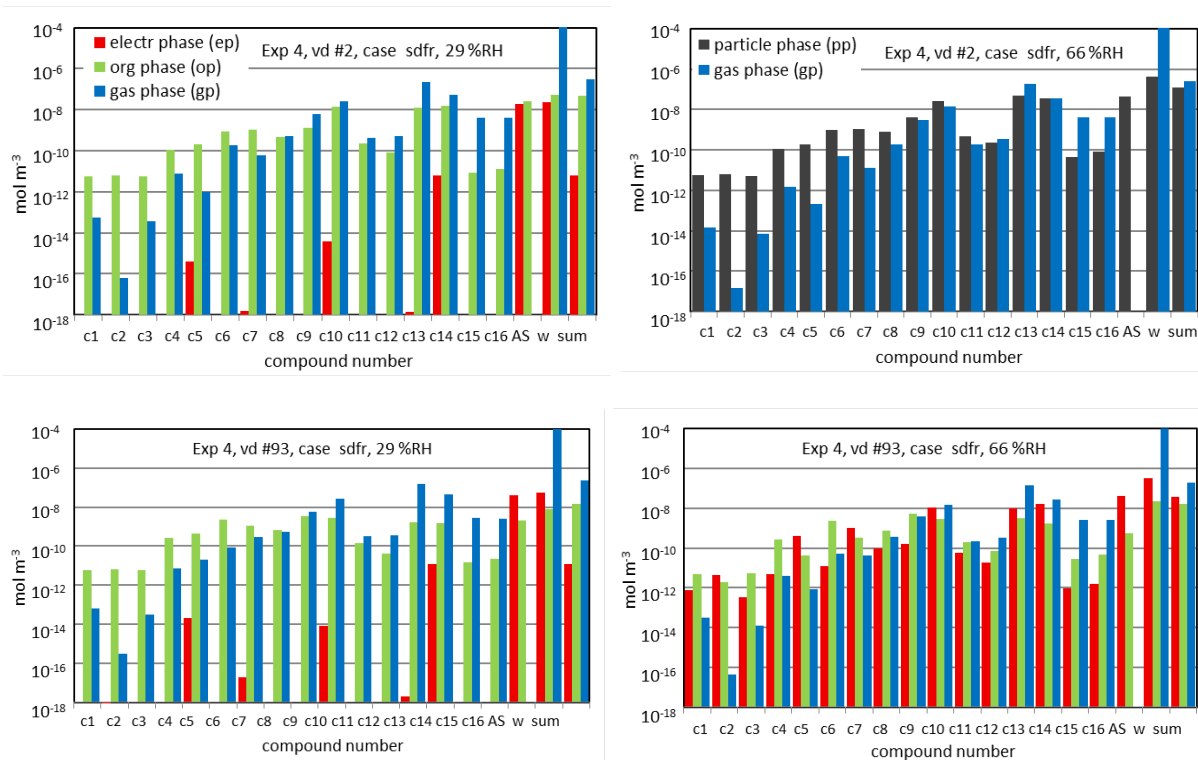


Figure 8. Calculated average O:C ratios for experiments at the measured RH (columns in full colors) and at the RH of the corresponding experiments (light colors, see Table 2 for reference) based on ten randomly chosen volatility distributions for the cases org, id, sd, sdfr, and orgfr (identified by different colors). The vertical black lines on the columns indicate the range of values obtained for the calculations with the individual volatility distributions. The horizontal dotted line marks the measured value for the experiment performed at low RH, the dashed line the value for the experiment performed at high RH.

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50 **Figure 9.** Equilibrium phase partitioning (mol m^{-3}) for case sdfr between gas phase, electrolyte phase, and organic phase at low RH (25 %) and at high RH (60 %) for experiment 8, volatility distribution #11 (upper panels) and experiment 13, volatility distribution #45 (lower panels). Compound numbers are c1: diaterpenylic acid acetate; c2: 3-MBTCA, c3: ValT4N9, c4: ValT4N10, c5: 3-hydroxyglutaric acid, c6: ValT4N3, c7: 3-oxoadipic acid; c8: pinic acid, c9: hopinonic acid, c10: glutaric acid; c11 norpinic acid, c12: 2-hydroxyterpenylic acid, c13: 5-COOH-3-OH-pentanal, c14: succinic acid, c15: 10-oxopinonic acid, c16: 4-oxopinonic acid, ammonium sulfate (AS), water (w), and the sum of compounds c1-c16.



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Figure 10. Equilibrium phase partitioning (mol m^{-3}) for case sdfr between gas phase, particle phase, electrolyte phase, and organic-rich phase at low (29 %) and high (66 %) RH for experiment 4, volatility distribution #2 (upper panels), and volatility distribution #93 (lower panels). Compound numbers are c1: diaterpenylic acid acetate; c2: 3-MBTCA, c3: ValT4N9, c4: ValT4N10, c5: 3-hydroxyglutaric acid, c6: ValT4N3, c7: 3-oxoadipic acid; c8: pinic acid, c9: hopinonic acid, c10: glutaric acid; c11 norpinic acid, c12: 2-hydroxyterpenylic acid, c13: 5-COOH-3-OH-pentanal, c14: succinic acid, c15: 10-oxopinonic acid, c16: 4-oxopinonic acid, ammonium sulfate (AS), water (w), and the sum of compounds c1-c16.

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