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Total Atmospheric Mercury Deposition in Forest Areas in Korea

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33 **Abstract**

34 Atmospheric mercury dry and wet deposition, mercury in throughfall and litterfall,
35 and mercury volatilization from soil were measured during August 2008 to February 2010 in
36 a temperate deciduous forest in Korea. The yearly estimated mercury budget was calculated
37 using two input approaches. For this location the annual mercury accumulation was estimated
38 to be $6.8 \mu\text{g m}^{-2} \text{yr}^{-1}$ or $3.9 \mu\text{g m}^{-2} \text{yr}^{-1}$ depending on the approach used. Cumulative wet and
39 throughfall fluxes were 4.3 and $6.7 \mu\text{g m}^{-2} \text{yr}^{-1}$, respectively. The annual litterfall flux was 4.6
40 $\mu\text{g m}^{-2} \text{yr}^{-1}$ and was highest from October to December due to the increased litter production
41 during that period. The annual Hg emission flux from soil was $6.8 \mu\text{g m}^{-2} \text{yr}^{-1}$. The overall
42 ratio of wet deposition, throughfall, and litterfall was 1: 1.6: 1.1. Cumulative dry deposition
43 fluxes of gaseous oxidized mercury (GOM) were highest in spring 2009 ($10.0 \pm 2.0 \mu\text{g m}^{-2}$
44 yr^{-1}), followed by summer 2009 ($5.8 \pm 4.2 \mu\text{g m}^{-2} \text{yr}^{-1}$), winter 2008 ($5.1 \pm 5.0 \mu\text{g m}^{-2} \text{yr}^{-1}$),
45 winter 2009 ($4.6 \pm 5.7 \mu\text{g m}^{-2} \text{yr}^{-1}$), fall 2008 ($1.9 \pm 1.0 \mu\text{g m}^{-2} \text{yr}^{-1}$) and fall 2009 (1.2 ± 1.4
46 $\mu\text{g m}^{-2} \text{yr}^{-1}$) while dry deposition fluxes for particulate bound mercury (PBM) were highest in
47 summer 2009 ($9.6 \pm 9.0 \mu\text{g m}^{-2} \text{yr}^{-1}$), followed by winter 2009 ($5.3 \pm 5.9 \mu\text{g m}^{-2} \text{yr}^{-1}$), winter
48 2008 ($3.8 \pm 2.0 \mu\text{g m}^{-2} \text{yr}^{-1}$), spring 2009 ($3.3 \pm 2.6 \mu\text{g m}^{-2} \text{yr}^{-1}$), fall 2008 ($3.0 \pm 1.7 \mu\text{g m}^{-2}$
49 yr^{-1}) and fall 2009 ($1.2 \pm 0.4 \mu\text{g m}^{-2} \text{yr}^{-1}$). The VWM TM concentration in throughfall ($14.4 \pm$
50 7.1 ng L^{-1}) was about two times higher than that in wet deposition ($5.9 \pm 3.8 \text{ ng L}^{-1}$). Wet
51 deposition and throughfall fluxes were higher in summer than those in other seasons possibly
52 due to a high precipitation depth.

53

54

55 **Keywords:** Dry deposition; Wet deposition; Throughfall; Litterfall; Gaseous oxidized
56 mercury (GOM); Particulate bound mercury (PBM); Hg emission flux

57



58 1. Introduction

59 Mercury (Hg) is a highly toxic pollutant and serious threat to human and ecosystem
60 health. It is classified as a persistent bioaccumulative and toxic (PBT) chemical (U.S.EPA,
61 1997a). Atmospheric Hg exists in three different forms with different chemical and physical
62 properties; gaseous elemental mercury (GEM, Hg^0), gaseous oxidized mercury (GOM, Hg^{2+}),
63 and particulate bound mercury (PBM, Hg_p). GEM is the major form of mercury in the
64 atmosphere and is relatively water insoluble and very stable with a long residence time of 0.5
65 - 2 year (Carpi and Lindberg, 1997; Cohen et al., 2004; Schroeder and Munthe, 1998).

66 GOM is water soluble, with relatively strong adhesion properties (Han et al., 2005)
67 and can be scavenged by rain within precipitating and below clouds (Lin and Pehkonen,
68 1999). It has very high dry deposition velocity similar to HNO_3 ($1\sim 5 \text{ cm sec}^{-1}$) if it is
69 assumed that all GOM is HgCl_2 (Petersen et al., 1995). PBM is strongly adsorbed to
70 atmospheric particulate matter including soot, dust, ice crystals, sea salts (Lu and Schroeder,
71 2004). Atmospheric PBM transport is significantly affected by its particle size distribution
72 and it is likely to be deposited at intermediate distances contributing to both wet and dry
73 deposition (Lynam and Keeler, 2002). Atmospheric deposition via wet and dry processes is
74 an important Hg input to the forest ecosystem (Buehler and Hites, 2002; Fitzgerald et al.,
75 1998; Landis and Keeler, 2002; Lindberg et al., 1998; Miller et al., 2005; Rolfhus et al.,
76 2003). Mercury deposited from the atmosphere can be transformed to methyl mercury
77 (MeHg) which bio-accumulates in aquatic food chains, resulting in adverse health and
78 ecological effects (Ma et al., 2013; Rolfhus et al., 2003). Therefore, monitoring of the
79 deposition flux as well as characterizing Hg deposition in forested areas is important for
80 estimating environmental risks associated with Hg.

81 The deposition processes of Hg in the forest ecosystem are very complicated because
82 of various interactions between atmospheric Hg and the canopy, such as oxidation of mercury
83 on leaf surfaces (Iverfeldt, 1991), deposition of GOM and PBM on leaf surfaces (St. Louis et
84 al., 2001), stomatal uptake of atmospheric Hg^0 (Iverfeldt, 1991; Lindberg et al., 1991; St.
85 Louis et al., 2001), uptake via roots of dissolved Hg in soil and soil water (Cocking et al.,
86 1995), and stomatal uptake of Hg^0 emitted from soils (Bishop et al., 1998).

87 Atmospheric Hg deposition to forests can follow several different pathways, such as
88 dry deposition, throughfall, and litterfall. Dry deposition is the deposition of pollutants,
89 including gases and particulate matter, onto surfaces including plant tissues. Hg deposited



90 onto plant surfaces can be revolatilized, incorporated into tissue or washed off by
91 precipitation (which is deemed throughfall) which often results in throughfall having higher
92 Hg concentrations than precipitation (Iverfeldt, 1991; Kolka et al., 1999; Munthe et al., 1995)
93 (Choi et al., 2008; Grigal et al., 2000; Schwesig and Matzner, 2000).

94 Litterfall is dead plant material such as leaves, bark, needles and twigs that has fallen
95 to the ground. Litterfall carries new Hg inputs from the atmosphere to the forest floor and
96 also Hg recycled from volatilization from soils and other surfaces. Throughfall and litterfall
97 contribute to the biochemical recycling of atmospheric mercury in forest systems (St. Louis et
98 al., 2001).

99 The objectives of this study were to characterize total atmospheric mercury
100 deposition in a temperate deciduous forested area in Korea by measuring mercury dry
101 deposition, wet deposition, throughfall, litterfall and volatilization from soils. Based on these
102 data, the annual mercury fluxes were estimated.

103

104

105 **2. Materials and methods**

106

107 **2.1. Site description**

108 The sampling sites were located at Yangsuri, Yangpyeong-gun, Gyeonggi-do, a
109 province in Korea where the Bukhan (North Han) and Namhan River (South Han River)
110 come together (**Fig. 1**). Gyeonggi-do has a population of 12 million (24% of the total
111 population and the most populated province in South Korea) and an area of 10,187 km² (10%
112 of the total area of South Korea). Yangpyeong-gun has a population of 83,000 (0.2% of the
113 total population in South Korea) and an area of 878.2 km² (0.9% of the total area in South
114 Korea). Wet deposition samples were collected at the Han River Environment Research
115 Center (Elevation 25 m, N37°32', E127°18') (site A in Fig. 1). Dry deposition, throughfall,
116 litterfall, volatilization from soils and TM in soil samples were determined in a deciduous
117 forest (Elevation 60 m, N37°32', E127°20') (site B in Fig. 1). Since this area contains rivers,
118 a flood plain, agricultural land, residential areas and forests, the study sites are appropriate for
119 identifying the transport and transformation of Hg, as well as in/out flow of Hg in a forested
120 ecosystem typical for this part of the world.



121

122 **2.2. Sampling methods**

123 Samples were collected from August 2008 to February 2010. Weekly samples for dry
124 and wet deposition in an open area and throughfall were collected using a dry and wet
125 deposition sampler (DWDS) in a deciduous forest dominated by chestnut.

126

127 2.2.1. Dry deposition for GOM and PBM

128 Similar to previous studies (Lai et al., 2011; Yi et al., 1996), the dry deposition
129 sampler was equipped with a knife-edge surrogate surface (KSS) sampler using quartz filter
130 for PBM deposition and KCl-coated quartz filter for GOM deposition. The quartz filter and
131 KCl-coated quartz filter (soaked in KCl solution for 12h and dried on clean bench) were pre-
132 baked in a quartz container at 900 °C. The filters were then placed on a filter holder base, and
133 clamped with a retaining ring before the filter holder was deployed in the KSS.

134

135 2.2.2. TM in wet deposition and throughfall

136 The DWDS for wet deposition and throughfall was equipped with four discrete
137 sampling systems that allows for two Hg and two trace elements sampling trains similar to
138 what was used in previous studies (Lai et al., 2007; Landis and Keeler, 1997; Seo et al., 2012;
139 Seo et al., 2015)

140

141

142 2.2.3. TM in soil and litterfall

143 Soil samples were collected every month from December 2008 to October 2010,
144 except January 2009, January, July, and August 2010, at a depths of 6 (A horizons) and 15 cm
145 (B horizons) in deciduous forest.

146 Litterfall samples was collected every month from December 2008 to November
147 2010, except January 2010. Ten nylon-mesh-lined baskets (1.09 m² each) were acid cleaned
148 and randomly placed under the canopy. All litter and soil samples were freeze-dried, sorted
149 by tree species, weighed, and then homogenized by crushing manually prior to analysis.

150

151 2.2.4. Volatilization from soils



152 The gaseous mercury emission flux from soil was measured using a dynamic flux
153 chamber (DFC) connected to the Tekran 2537A (Tekran Inc., Toronto, Canada) and Tekran
154 1110 dual sampling unit (allows alternate sampling from inlet and outlet) (Choi and Holsen,
155 2009) under the deciduous forest area once a month. Daily automated calibrations were
156 performed for the Tekran 2337A using an internal permeation source. Manual injections were
157 used to evaluate these calibrations using a saturated mercury vapor standard. The flowrate
158 was approximately 5 L min⁻¹. Four 1 cm diameter inlet holes were evenly placed around the
159 chamber ensuring it was well mixed. The DFC (3.78L) were placed 2cm under the soil. The
160 types of DFCs were glass and polycarbonate which may block some UV light.

161

162 **2.3. Analytical method**

163

164 2.3.1. Dry deposition for GOM and PBM

165 The dry deposition samples for GOM and PBM samples were analyzed using a tube
166 furnace connected to a Tekran 2537. The tube furnace was pre-heated (GOM: 525 °C, PBM:
167 900 °C) and zero air passed through until the Hg concentration was zero. After samples were
168 placed inside the tube furnace, the tube furnace was purged with zero air until Hg level was
169 again zero. The mass of Hg desorbed from the sample was determined using the product of
170 concentration and flowrate (5 L min⁻¹). The system recovery was measured by injecting
171 mercury vapor standards (0, 10, 20, 30, 50 µL) manually.

172

173 2.3.2. TM in wet deposition and throughfall

174 Total mercury (TM) in throughfall was measure using a Tekran Series 2600 equipped
175 with cold vapor atomic fluorescence spectrometer (CVAFS) following the procedures
176 outlined in the U.S. EPA Method 1631 version E (U.S.EPA, 2002) and the U.S. EPA Lake
177 Michigan Mass Balance Methods Compendium (LMMBMC) (U.S.EPA, 1997b).

178

179 2.3.3. TM in soil and litterfall

180 Total mercury concentrations in soil and litterfall samples were determined using a
181 direct mercury analyzer (DMA-80, Milestone, Italy), which utilizes the serial process of
182 thermal composition, catalytic reduction, amalgamation, desorption, and atomic absorption
183 spectroscopy.



184

185 **2.4. QA/QC**

186 2.4.1 Dry deposition for GOM and PBM

187 Relative percent difference (RPD) analyses for replicate GOM and PBM

188 measurements were 19.4% and 22.9%, respectively. Recovery (%) of the Tekran 2537

189 measured by direct injection was 87~107% ($r^2 > 0.9995$). The MDL was 0.04 ng L^{-1} , the

190 same as reported in EPA Method 1631. Relative Standard Deviation (RSD) measured by

191 injecting mercury vapor standards in the same concentration seven times averaged 2 ~ 5%,

192 within EPA Method 1631 requirements ($\pm 25\%$). The average field blank ($n = 4$)

193 concentration was 0.36 ng L^{-1} and the average lab blank ($n = 44$) concentration was 0.2 ng L^{-1} .

194 ¹.

195

196 2.4.2. TM in wet deposition and throughfall

197 The method detection limit (MDL) for TM in wet deposition and throughfall

198 calculated as three times the standard deviation of seven sequential reagent blanks was 0.05 ng

199 L^{-1} . Initial Precision Recovery (IPR) and On-going Precision Recovery (OPR) measured

200 every 15 samples at the start of the analysis ranged from 80 ~ 107% ($92.2 \pm 7.0 \%$ in average)

201 and 81 ~ 117% ($96.9 \pm 13.7 \%$ in average), respectively. The analysis of matrix spikes (MS)

202 and matrix spike duplicates (MSD) measured every 10 samples to assess accuracy and

203 precision ranged from 80 ~ 123% and relative percent difference (RPD) was 3 ~ 13%.

204

205 2.4.3. TM in litterfall and soil

206 TM in litterfall and soil was reported on a dry-weight basis. Recovery (%) of

207 standard reference materials (SRMs) (MESS3, marine sediment) purchased from the National

208 Research Council of Canada and analyzed every 10 samples at the start of experiments was

209 $104 \pm 4\%$.

210

211 2.4.4. Volatilization from soil

212 The DFC was connected to the Tekran 2537A through Tekran 1110 sampling unit.

213 Ten μL of vapor phase Hg was injected into the DFC ($n = 10$) before deployment in the field.

214 Recovery was 86 ~ 110% and averaged 101% at a flow rate of 5 L min^{-1} .

215



216

217 **3. Results and Discussion**

218

219 **3.1. Monthly and seasonal variations in dry deposition fluxes of GOM and PBM**

220 The average dry deposition fluxes for GOM and PBM were $5.4 \mu\text{g m}^{-2} \text{yr}^{-1}$ (range:
221 $0.4 \sim 14.4 \mu\text{g m}^{-2} \text{yr}^{-1}$) and $0.5 \text{ ng m}^{-2} \text{h}^{-1}$ (range: $0.8 \sim 19.4 \mu\text{g m}^{-2} \text{yr}^{-1}$), respectively (Fig.
222 S1). The dry deposition fluxes for GOM were highest in spring 2009 ($10.0 \pm 2.0 \mu\text{g m}^{-2} \text{yr}^{-1}$),
223 followed by summer 2009 ($5.8 \pm 4.2 \mu\text{g m}^{-2} \text{yr}^{-1}$), winter 2008 ($5.1 \pm 5.0 \mu\text{g m}^{-2} \text{yr}^{-1}$), winter
224 2009 ($4.6 \pm 5.7 \mu\text{g m}^{-2} \text{yr}^{-1}$), fall 2008 ($1.9 \pm 1.0 \mu\text{g m}^{-2} \text{yr}^{-1}$) and fall 2009 ($1.2 \pm 1.4 \mu\text{g m}^{-2}$
225 yr^{-1}) while the dry deposition fluxes for PBM were highest in summer 2009 ($9.6 \pm 9.0 \mu\text{g m}^{-2}$
226 yr^{-1}), followed by winter 2009 ($5.3 \pm 5.9 \mu\text{g m}^{-2} \text{yr}^{-1}$), winter 2008 ($3.8 \pm 2.0 \mu\text{g m}^{-2} \text{yr}^{-1}$),
227 spring 2009 ($3.3 \pm 2.6 \mu\text{g m}^{-2} \text{yr}^{-1}$), fall 2008 ($3.0 \pm 1.7 \mu\text{g m}^{-2} \text{yr}^{-1}$) and fall 2009 (1.2 ± 0.4
228 $\mu\text{g m}^{-2} \text{yr}^{-1}$) (Fig. 2). Nonparametric Mann-Whitney tests indicated that there were
229 statistically significant differences in the dry deposition fluxes for GOM between spring
230 2009, fall 2008, and fall 2009 ($p < 0.05$) and there were statistically significant differences in
231 the dry deposition flux for PBM between summer 2009 and fall 2009 ($p < 0.05$).

232 Zhang et al. (2012) reported that in eastern and central North America the GEM
233 concentration in the colder season were generally higher than in warmer seasons. However,
234 the dry deposition fluxes for GOM and PBM in spring and summer (warmer seasons) were
235 higher than in the fall and winter (cold seasons) following the same pattern as average GEM
236 concentrations (summer 2009: $2.7 \pm 0.9 \text{ ng m}^{-3}$, spring 2009: $2.4 \pm 0.6 \text{ ng m}^{-3}$, fall 2009: 2.3
237 $\pm 0.7 \text{ ng m}^{-3}$, winter 2008: $1.2 \pm 0.2 \text{ ng m}^{-3}$) in Han River Environment Research Center
238 (located approximately 2 km away). This suggests that GEM contributes to the measured dry
239 deposition (GOM, PBM) (Zhang et al., 2012).

240 Previous studies reported that the Hg species were dry deposited to leaf surfaces
241 through oxidation, adsorption (Munthe et al., 1995) and uptake by stomata (Lindberg et al.,
242 1992). As will be discussed later, the observed Hg dry deposition fluxes were also compared
243 with the estimated Hg dry deposition fluxes using litterfall, throughfall and wet deposition.

244

245

246 **3.2. Monthly and seasonal variations of TM wet deposition and throughfall flux**



247 The average VWM TM concentration in precipitation ($n = 35$) was highest in winter
248 2009 ($7.8 \pm 4.6 \text{ ng L}^{-1}$) ($n = 3$), followed by fall 2009 ($6.7 \pm 2.6 \text{ ng L}^{-1}$) ($n = 8$), winter 2008
249 ($6.3 \pm 5.7 \text{ ng L}^{-1}$) ($n = 3$), fall 2008 ($5.8 \pm 3.9 \text{ ng L}^{-1}$) ($n = 5$), spring 2009 ($5.0 \pm 3.5 \text{ ng L}^{-1}$) (n
250 $= 6$), and summer 2009 ($4.0 \pm 2.5 \text{ ng L}^{-1}$) ($n = 10$) (Fig. 3). Nonparametric Mann-Whitney
251 tests indicated that there were no statistically significant differences in the VWM TM
252 concentration between winter 2009 and other seasons which is probably related with the
253 small number of samples.

254 The average VWM TM concentration in throughfall ($n = 44$) was also highest in
255 winter 2009 ($32.4 \pm 6.7 \text{ ng L}^{-1}$) ($n = 7$), followed by winter 2008 ($21.6 \pm 17.8 \text{ ng L}^{-1}$) ($n = 3$),
256 fall 2008 ($10.1 \pm 6.1 \text{ ng L}^{-1}$) ($n = 5$), fall 2009 ($9.1 \pm 2.7 \text{ ng L}^{-1}$) ($n = 9$), spring 2009 ($8.5 \pm$
257 5.1 ng L^{-1}) ($n = 7$), and summer 2009 ($4.9 \pm 4.5 \text{ ng L}^{-1}$) ($n = 13$). VWM TM concentration in
258 winter 2009 was statistically significantly higher than fall 2009 ($p = 0.007$), spring 2009 ($p =$
259 0.035), and summer 2009 ($p = 0.001$).

260 The high VWM TM concentrations in precipitation and throughfall in winter were
261 associated with the combined effects of reduced mixing heights (Kim et al., 2009; Seo et al.,
262 2015), and low rainfall depth (11.7% of total rainfall depth) which is a typical pattern in
263 Yangpyung, Korea (KMA,
264 http://www.kma.go.kr/weather/climate/average_30years.jsp?yy_st&tnqh_x003D;2011&stn&tnqh_x003D;108&norm&tnqh_x003D;M&obs&tnqh_x003D;0&mm&tnqh_x003D;5&dd&tnqh_x003D;25&x&tnqh_x003D;25&y&tnqh_x003D;5
265 (accessed November 24, 2015)). Another possible reason for the high TM concentration in
266 precipitation and throughfall in winter was due to snow events. Scavenging by snow is more
267 efficient than by rain due to the larger surface area of snow (snow: $700 \text{ cm}^2/\text{g}$, rain: $60 \text{ cm}^2/\text{g}$)
268 (Kerbrat et al., 2008).
269
270

271 Previous studies reported that rainfall depth in forested areas were approximately
272 8~24% smaller than that in an open area (Choi et al., 2008; Deguchi et al., 2006; Keim et al.,
273 2005; Price and Carlyle-Moses, 2003) due to capture by the foliage and subsequent
274 evaporation. In this study, rainfall depth in the forest was approximately 8% smaller than that
275 in an open area. The TM concentration in throughfall was higher than in precipitation
276 (statistically significant differences ($r^2 = 0.20$) ($p < 0.05$)) due to wash off of previously
277 deposited Hg from the foliage (Grigal et al., 2000; Iverfeldt, 1991; Kolka et al., 1999;
278 Schwesig and Matzner, 2000) and oxidation of Hg^0 to Hg^{2+} on the wet foliage surface by



279 ozone and subsequent wash off (Graydon et al., 2008). Other possible sources of Hg in
280 throughfall are leaching and mechanical weathering of Hg from foliage. Some of the
281 deposited Hg can be washed off by rainfall and reemitted to the atmosphere (Rea et al.,
282 2001). Therefore, all of the Hg deposited on the foliar surfaces is not in the throughfall.

283
284

285 **3.3 Relationship between rainfall depth, VWM TM concentration, TM wet deposition** 286 **and throughfall flux**

287

288 There was a statistically significant negative correlation between rainfall depth and
289 VWM TM concentrations in precipitation ($r^2 = 0.13$) ($p < 0.05$) (Fig. S2) and throughfall ($r^2 =$
290 0.19) ($p < 0.05$) (Fig. S3) due to dilution during the later stage of a precipitation event. This
291 negative correlation has been also found in previous studies (Guo et al., 2008; Landis and
292 Keeler, 2002; Seo et al., 2012; Seo et al., 2015; Wallschläger et al., 2000). About 19% of
293 throughfall and 13% of precipitation variation in VWM concentration are explained by
294 precipitation depth, the rest of the variation is likely due to variations in local and regional
295 sources and other mechanisms involved in biogeochemical cycling for example dry
296 deposition (St. Louis et al., 2001). There was a statistically significant positive correlation
297 between rainfall depth and TM deposition flux in precipitation ($r^2 = 0.34$) ($p < 0.05$),
298 suggesting that the TM deposition flux increased during large events even though continuous
299 rain diluted the TM mass, similar to previous studies (Choi et al., 2008; Wang et al., 2014).
300 However, a large rainfall depth does not affect wet deposition fluxes if atmospheric
301 concentrations of GOM and PBM are low (Zhang et al., 2012).

302

303 **3.4. Leaf-on vs. Leaf-off**

304 At this sampling site the leaf-on season is from March to the end of November.
305 During leaf-on periods, the TM concentrations in throughfall (deciduous trees) (average 8.1
306 ng L^{-1}) were higher than that in precipitation (average 5.4 ng L^{-1}) and they were significantly
307 correlated ($r^2 = 0.59$) ($p < 0.05$). For leaf-off periods TM concentrations in throughfall
308 (average 14.3 ng L^{-1}) were 1.7 times higher than in precipitation (average 8.6 ng L^{-1}) and
309 concentrations were moderately correlated ($r^2 = 0.56$) ($p < 0.05$) (Table 1). The concentration
310 enhancement during leaf-off periods was probably due, at least in part, to snow on the



311 branches that collected mercury due to dry deposition during dry periods that was
312 subsequently collected by the sampler after being blown off by wind or after it melted.

313 The flux of Hg in throughfall was similar to or lower than that of precipitation
314 although the TM concentration in throughfall was higher than that in precipitation. The
315 cumulative Hg fluxes in throughfall (leaf on: $7.0 \mu\text{g Hg m}^{-2}$, leaf off: $3.1 \mu\text{g Hg m}^{-2}$) were
316 higher than in precipitation (leaf on: $4.9 \mu\text{g Hg m}^{-2}$, leaf off: $0.6 \mu\text{g Hg m}^{-2}$). As mentioned
317 previously this may be a result of differences in rainfall depth (leaf-on periods) and snow
318 events (leaf-off periods).

319

320

321 **3.5. TM in litterfall and soil**

322 Litterfall can be an important Hg input to soils under forested landscapes. The mean
323 monthly TM concentrations in litterfall were $50.2 \pm 16.5 \text{ ng g}^{-1}$ for the deciduous forest
324 (ranged from 28.2 to 76.4 ng g^{-1}) (Fig. 4). TM litterfall fluxes from winter 2009 to fall 2010
325 (one year) were $0.3 \pm 0.4 \mu\text{g m}^{-2}$ in the deciduous forest (ranged from 0.01 to $1.9 \mu\text{g m}^{-2}$). TM
326 litterfall fluxes were different depending on the sampling periods; being lowest in summer,
327 from June to August, and highest in fall, from September to November (Fig. 4) because
328 litterfall production increases substantially over the growing season, from late fall to early
329 winter. Hall and St. Louis (2004) reported the mean concentration of TM in leaf litter
330 increased from 7.1 ng g^{-1} to a final value of 40.9 ng g^{-1} in deciduous litter. Demers et al.
331 (2007) reported that the quantity of TM added to the decaying deciduous leaf litter was $5.1 \sim$
332 $5.5 \mu\text{g m}^{-2}$, during the growing season. In this study, TM litterfall fluxes were smaller than
333 those in previous studies.

334 Soil samples were collected from the near-surface A-horizon following the removal
335 of any rock fragments and B-horizon. The mean soil TM concentrations were higher within
336 the A-horizon ($66.9 \pm 20.3 \text{ ng g}^{-1}$) than within the B-horizon ($46.1 \pm 17.5 \text{ ng g}^{-1}$) deciduous
337 forest stand. TM concentration in soil collected in this study was similar to TM concentration
338 found in soil collected from uncontaminated baseline sites which ranged from 30 to 50 ng g^{-1}
339 (Gray et al., 2015).

340

341 **3.6. Volatilization from soils**



342 Hg emission fluxes were estimated from directly measured soil volatilization of
343 gaseous elemental mercury (GEM) using a dynamic flux chamber (DFC). The measured
344 fluxes were the highest in June, the lowest in November. Emission fluxes were positively
345 correlated with ambient air temperature however, they were not influenced by precipitation.
346 For example the ambient air temperature was higher in summer than other seasons, but were
347 not higher in July, a period of several severe rain storms nor were they lower in August which
348 had very little rain. This result may be because the relative humidity was high enough that the
349 soil remained moist. This result is similar to a previous study that found that Hg emission
350 fluxes were positively correlated with soil surface temperature and negatively correlated with
351 humidity (Choi and Holsen, 2009; Gabriel et al., 2006; Wallschläger et al., 2000; Wang et al.,
352 2005). Hg emission fluxes during leaf-on periods (March to November) ($0.65 \pm 2.25 \text{ ng m}^{-2}$
353 hr^{-1} , $16.9 \text{ }^\circ\text{C}$) were higher than leaf-off periods (December) ($0.02 \pm 2.13 \text{ ng m}^{-2} \text{ hr}^{-1}$, $-1.29 \text{ }^\circ\text{C}$).
354 This result is similar to previous study. Choi and Holsen (2009) reported that during leaf-off
355 periods, Hg emission flux were correlated with temperature and solar radiation. The
356 cumulative annual Hg emission fluxes was $6.8 \text{ } \mu\text{g m}^{-2} \text{ yr}^{-1}$ (Fig. 5). Due to sampler (Tekran
357 2537A) malfunctions in January, February and April fluxes were assumed to be equal to the
358 average of the flux the month before and after.

359

360 3.7 Estimated dry deposition at forest

361 Fu et al. (2009) estimated dry deposition to be equal to litterfall + throughfall – wet
362 deposition. Using the data presented here, the estimated dry deposition flux ($6.7 \text{ } \mu\text{g m}^{-2} \text{ yr}^{-1}$)
363 was lower than measured dry deposition ($9.9 \text{ } \mu\text{g m}^{-2} \text{ yr}^{-1}$) and there was no significant
364 correlation between the two methods ($r^2 = 0.22$) ($p = 0.65$). The differences in the estimates
365 could be due to the loss of litter samples by wind or Hg losses from the collected litter due to
366 meteorological conditions such as rainfall (Blackwell et al., 2014) due to relatively long
367 sampling periods (1 month). However dry deposition collected with a surrogate surface
368 doesn't include accumulation in leaf stomata which may underestimate dry deposition using
369 this technique. Another reasons for these differences could be deposition to all media such as
370 leaves, tree branch, soils and land types surrounding the sites (Zhang et al., 2012).

371 The annual input flux calculated by summing wet deposition plus dry deposition
372 ($14.3 \text{ } \mu\text{g m}^{-2} \text{ yr}^{-1}$) was higher than the input flux calculated by summing throughfall +
373 litterfall ($12.8 \text{ } \mu\text{g m}^{-2} \text{ yr}^{-1}$) (Fig. 6). Nonparametric Mann-Whitney tests indicated that there



374 were not statistically significant differences ($r^2 = 0.14$) ($p = 0.98$). In general wet + dry
375 deposition was larger than throughfall plus litterfall except during fall when leaves were
376 being actively dropped from the trees. The largest difference was in July during a period of
377 significant precipitation (about 26.3 % of the total amount in 2009). This difference is most
378 likely due to the many reactions and transformations on the leaf surface that aren't mimicked
379 with the surrogate surface including re-emission (Rea et al., 2001).

380

381 **3.8. Mercury budget**

382 The yearly estimated mass balance of mercury was calculated using both input
383 approaches described above. Input to the forest canopy (wet deposition in an open area: 4.3
384 $\mu\text{g m}^{-2} \text{yr}^{-1}$, dry deposition in the forested area: 9.9 $\mu\text{g m}^{-2} \text{yr}^{-1}$) with output (emissions from
385 soil 6.8 $\mu\text{g m}^{-2} \text{yr}^{-1}$, TM in soil: 0.6 $\mu\text{g m}^{-2} \text{yr}^{-1}$) was estimated to be 6.8 $\mu\text{g m}^{-2} \text{yr}^{-1}$. The
386 alternative method with input (throughfall: 6.7 $\mu\text{g m}^{-2} \text{yr}^{-1}$, litterfall: 4.6 $\mu\text{g m}^{-2} \text{yr}^{-1}$) with
387 output (emissions from soil 6.8 $\mu\text{g m}^{-2} \text{yr}^{-1}$, TM in soil: 0.6 $\mu\text{g m}^{-2} \text{yr}^{-1}$) results in a net
388 mercury flux of 3.9 $\mu\text{g m}^{-2} \text{yr}^{-1}$.

389 This result is similar to previous studies. At the Lehstenbach catchment in Germany
390 the estimated fluxes was 6.8 $\mu\text{g m}^{-2} \text{yr}^{-1}$ (Schwesig and Matzner, 2000). In the Experimental
391 Lakes Area (ELA) watersheds in Canada the flux was 3 ~4 $\mu\text{g m}^{-2} \text{yr}^{-1}$ (St. Louis et al., 2001).
392 However, for the Lake Langtjern spruce forest in southeast Norway (20.1 $\mu\text{g m}^{-2} \text{yr}^{-1}$)
393 (Larssen et al., 2008) and Huntington Wildlife forest (15.9 $\mu\text{g m}^{-2} \text{yr}^{-1}$ in deciduous, 26.8 μg
394 $\text{m}^{-2} \text{yr}^{-1}$ in conifer) (Blackwell et al., 2014) the estimated fluxes were higher than in this
395 study.

396 There are several few uncertainties associated with this study. Dry deposition
397 measured with the surrogate surface does not account for accumulation in leaf stomata yet
398 this technique yielded a larger flux than to litterfall + throughfall – wet deposition.

399 Litterfall can be lost from the sampler by wind or and Hg can be lost from the
400 collected litter due to rainfall due to relatively long sampling periods (Blackwell et al., 2014)
401 and approximately half of the GEM stored in the leaf may be released to back to the
402 atmosphere (Zhang et al., 2012).

403 DFCs can alter fluxes because they cover the soil potentially blocking some UV
404 light. In addition several months of measurements were missed. Grab samples for TM in soil



405 may not capture the true variability in the forest. Therefore, further investigations are
406 required to elucidate exactly the mercury budget Additional work should focus on better
407 quantifying dry deposition, TM in soil water, overflow rate and biogeochemically recycling
408 within forest.

409
410

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416

417 **References**

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561

562 **Table List**

563 Table 1. Cumulative precipitation depths, VWM Hg concentration, cumulative Hg fluxes in
564 precipitation and throughfall during leaf-on and leaf-off periods.

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567

568 **Figure List**

569 Fig. 1. The locations of the sampling sites used in this study (Yangsu-ri, Korea)

570 Fig. 2. Seasonal variation in dry deposition flux for GOM and PBM under the deciduous
571 forest.

572 Fig. 3. Seasonal variation in VWM TM concentration, rainfall depth and TM flux in
573 precipitation and throughfall.

574 Fig. 4. Seasonal variation in TM concentration and flux in a deciduous forest.

575 Fig. 5. The estimated annual Hg emission fluxes in 2009 from soil.

576 Fig. 6. Comparison of Deposition flux calculated by summing wet deposition + dry
577 deposition and throughfall + litterfall

578



579 Table 1. Cumulative precipitation depths, VWM Hg concentration, cumulative Hg fluxes in
580 precipitation and throughfall during leaf-on and leaf-off periods.

	Cumulative precipitation depth (mm)		VWM Hg Concentration (ng L ⁻¹)		Cumulative Hg fluxes (µg Hg m ⁻²)	
	Leaf-on	Leaf-off	Leaf-on	Leaf-off	Leaf-on	Leaf-off
Precipitation	968.3	117.6	5.4	7.2	3.8	0.5
Throughfall	1009.7	114.7	8.1	18.3	4.9	1.8

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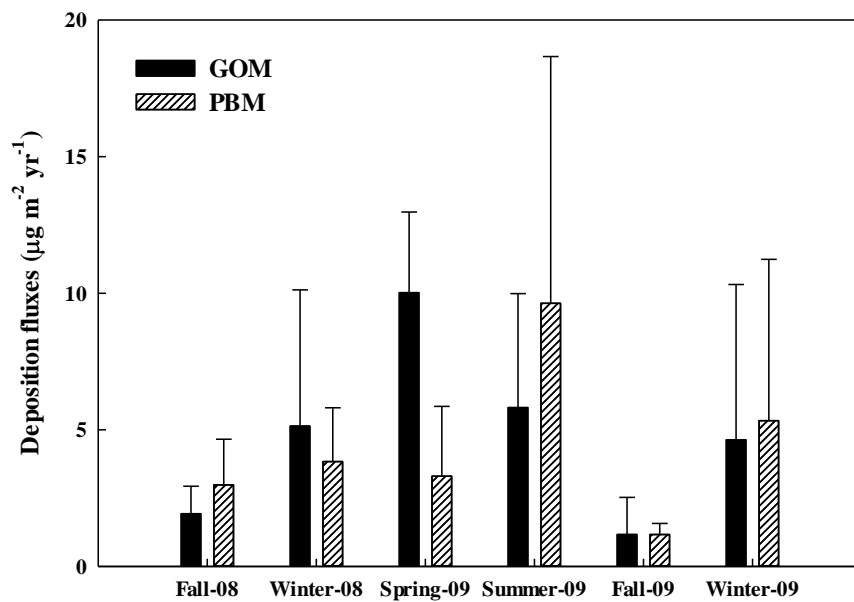


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Fig. 1. The locations of the sampling sites used in this study (Yangsu-ri, Korea).

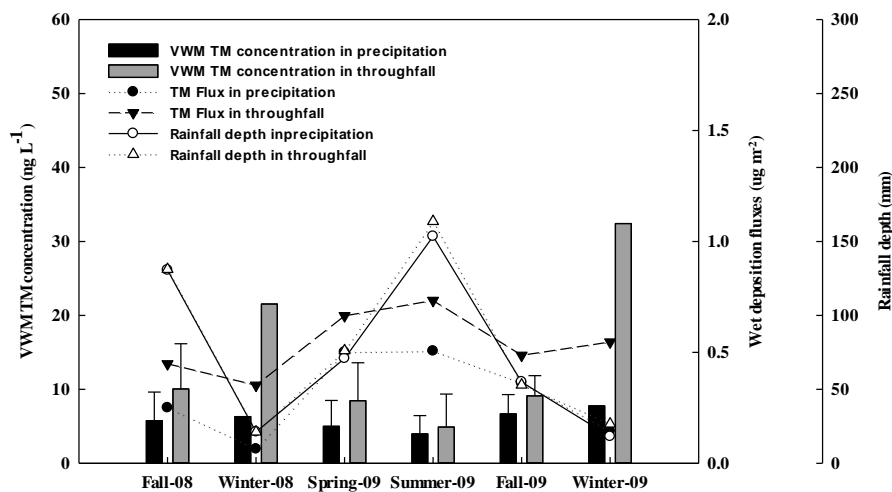
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587 **Fig. 2. Seasonal variation in dry deposition flux for GOM and PBM under the**
588 **deciduous forest.**

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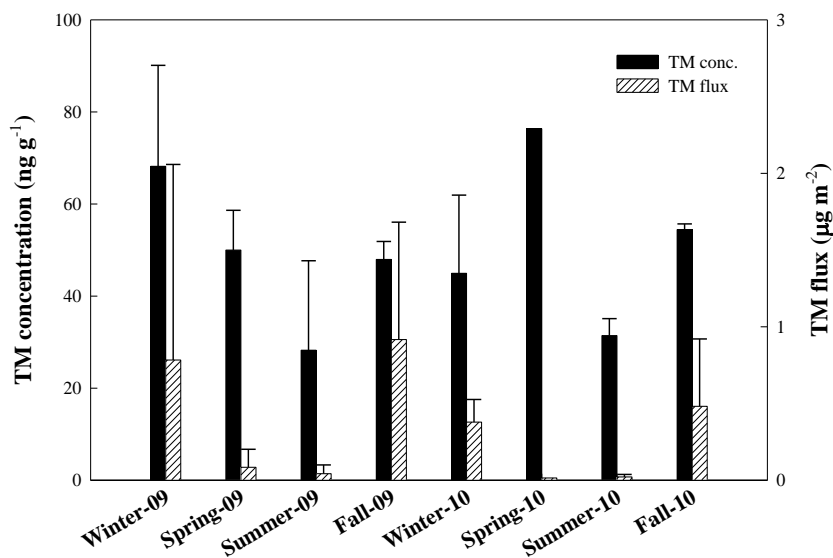
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592 **Fig. 3. Seasonal variation in VWM TM concentration, rainfall depth and TM flux in**
593 **precipitation and throughfall.**

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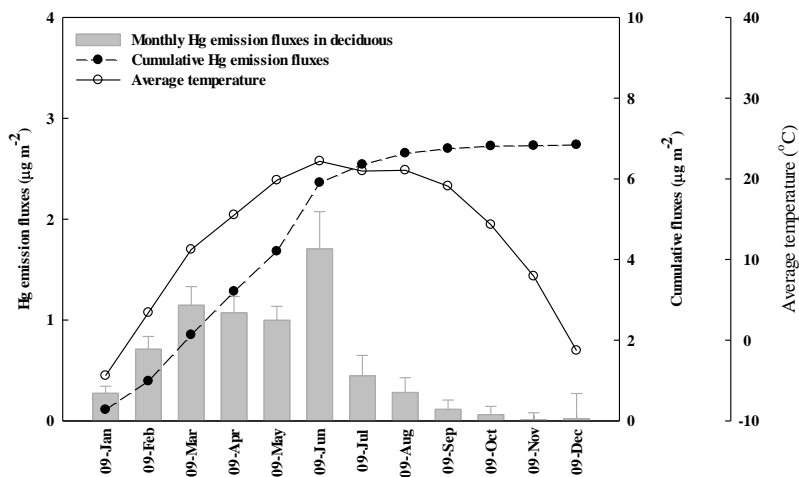
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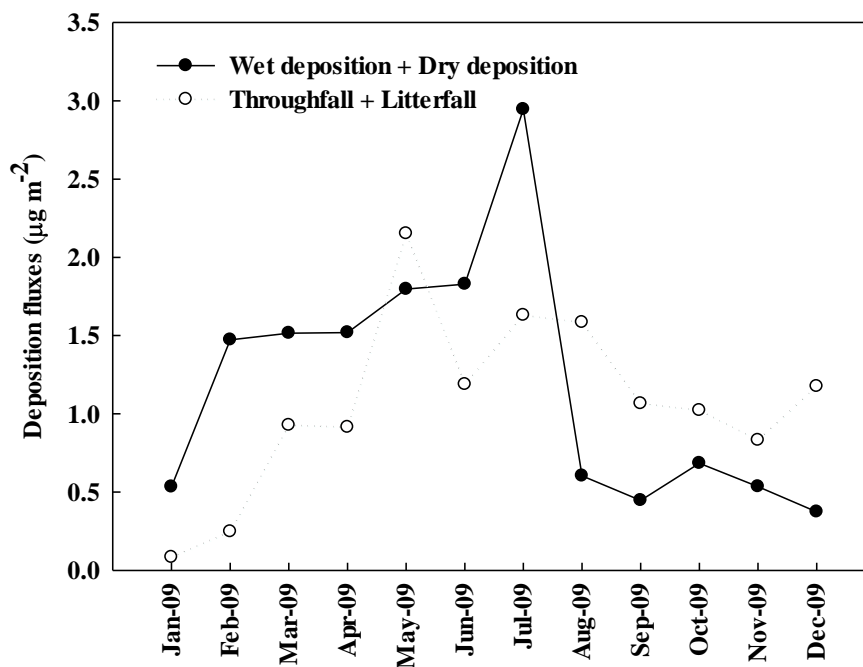
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Fig. 4. Seasonal variation in TM concentration and flux in a deciduous.



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Fig. 5. The estimated annual Hg emission fluxes in 2009 from soil.



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610 **Fig. 6. Comparison of Deposition flux calculated by summing wet deposition + dry**
611 **deposition and throughfall + litterfall**