Atmospheric aerosols in Rome, Italy: Sources, dynamics and spatial variations during two seasons

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Abstract

Investigations on atmospheric aerosols and their sources were performed during October/November 2013 and May/June 2014 subsequently in a suburban area of Rome (Tor Vergata) and in central Rome (near St. Peter's Basilica). During both years a Saharan dust advection event temporarily increased PM_{10} concentrations at ground level by some 12-17 µg m^{-3}. Generally, during Oct/Nov the ambient aerosol was more strongly influenced by primary emissions, whereas higher relative contributions of secondary particles (sulphate, aged organic aerosol) were found during May/June. Absolute concentrations of anthropogenic emission tracers (e.g. NO_x, CO_2, particulate polyaromatic hydrocarbons, traffic-related organic aerosol) were generally higher at the urban location. Positive matrix factorisation was applied to the PM_1 organic aerosol (OA) fraction of aerosol mass spectrometer (HR-ToF-AMS) data in order to identify different sources of primary OA (POA): traffic, cooking, biomass burning, and (local) cigarette smoking. While biomass burning OA was only found at the suburban site, where it accounted for the major fraction of POA (18-24 % of total OA), traffic and cooking were more dominant sources at the urban site. A particle type associated with cigarette smoke emissions, which is associated with a potential characteristic marker peak (*m/z* 84, $C_5H_{10}N^+$, a nicotine fragment) in the mass spectrum, was only found in central Rome, where it was emitted in close vicinity to the measurement location. Regarding secondary OA, in Oct/Nov, only a very aged, regionally advected oxygenated OA was found, which contributed 42-53 % to the total OA. In May/June total oxygenated OA accounted for 56–76 % of the OA. Here also a fraction (18-26 % of total OA) of a fresher, less oxygenated OA of more local origin was observed. New particle formation events were identified from measured particle number concentrations and size distributions during May/June 2014 at both sites. While they were observed every day at the urban location, at the suburban location they were only found under favourable meteorological conditions, but independent of advection of the Rome emission plume. Particles from sources in the metropolitan area of Rome and particles advected from outside Rome contributed 42-70 % and 30-58 % to total measured PM_1, respectively. Apart from the general aerosol characteristics, in this study the properties (e.g. emission strength) and dynamics (e.g. temporal behaviour) of each identified aerosol type is investigated in detail in order to provide a better understanding of the observed seasonal and spatial differences.

1 Introduction

Atmospheric aerosol particles remain a major uncertainty in both, estimations of climate change (Boucher et al., 2013) and of impact of air pollution on public health (Heal et al., 2012), and therefore are a major topic of current research (Fuzzi et al., 2015). Identifying the sources, properties and concentrations of atmospheric particles is essential for evaluating their effect on climate and health and constitutes a crucial step in finding measures for the improvement of air quality.

Many studies on aerosols and their sources have been performed in urban environments (e.g. Freutel et al., 2013; Mohr et al., 2012; Zheng et al., 2005), which are characterized by high population densities and a large diversity of particle sources. Typical urban aerosol sources include road traffic, cooking, and heating activities. Also emissions from biomass burning can be important, both of regional origin (e.g. agricultural and wild fires; Reche et al., 2012), and from residential wood combustion, which recently has become more prominent in Europe even in urban environments (Fuller et al., 2013).

Many of these anthropogenic sources emit large amounts of organic material in the fine particle fraction (e.g. Hildemann et al., 1991). In recent studies of particle source identification (e.g. Allan et al., 2010; Mohr et al., 2012; Reche et al., 2012), positive matrix factorisation (PMF) was applied to separate the organic aerosol (OA) fraction into different factors associated with various OA sources, thereby providing indications about the fraction of primary and secondary organic aerosol (POA and SOA) (Zhang et al., 2011). Oxygenated organic aerosol (OOA), mainly associated with SOA, is typically found to be the most abundant fraction of OA (Lanz et al., 2010), with concentrations depending on season and location (Zhang et al., 2011). Several studies, mainly such from observations during summer time (Lanz et al., 2010), show discrimination of OOA into a fresher and a more aged type of OOA based on different states of oxygenation and/or volatility (Jimenez et al., 2009).

While AMS measurements yield useful information on the age of OA, they cannot provide evidence for new particle formation of fresh secondary aerosol. Indications for such, however, can be found in physical aerosol properties like particle number concentration or size distributions (e.g. Alam et al., 2003). New particle formation events in urban environments have been investigated previously in several studies (e.g. Alam et al., 2003; Brines et al., 2015; Minguillon et al., 2015; Shi et al., 2001; Zhang et al., 2004), and especially in the early afternoon seem to be responsible for elevated particle number concentrations in urban areas in Southern Europe (Reche et al., 2011).

On the other hand, while the health impact of coarse particles (PM_{10}-$PM_{2.5}$) is not yet fully understood (Heal et al., 2012), the association between Saharan dust advections and mortality/hospitalisation is quite well demonstrated (Stafoggia et al., 2016). Deserts are large sources for mineral dust, which can strongly contribute to atmospheric PM_{10} levels, especially in Southern Europe. Measurements performed in the period 2001-2004 during Saharan dust advections over Rome showed a mean Saharan dust contribution of 12-16 $\mu g\ m^{-3}$ to daily PM_{10} concentrations, leading to an average annual increase of about 2 $\mu g\ m^{-3}$ (Gobbi et al., 2013). In the central Mediterranean region, maximum dust concentrations are typically observed from spring to autumn (Barnaba and Gobbi, 2004).

In this study, we investigate the occurrence and properties of ambient aerosol from different types of sources in Rome, which apart from local emissions can be influenced by advected aerosol from continental Europe and the Sahara desert. During two different seasons (Oct/Nov 2013 and May/June 2014) and at two different locations (city centre and suburb), stationary measurements of chemical and physical properties of aerosols, several trace gases, and meteorological variables were performed. Non-refractory components of submicron particles were measured with an Aerodyne high-resolution time-of-flight aerosol mass spectrometer (HR-ToF-AMS). To support identification of particle sources, their strength and temporal behaviour, the OA measured with the HR-ToF-AMS was further separated into different factors using PMF.

Based on these measurements, in this work the urban atmosphere of Rome is investigated in terms of particle source identification with a special focus on seasonal and spatial differences influencing the presence and/or the characteristics of aerosol types in the city area.

2 Experimental

2.1 Measurement locations and periods

Measurement results presented in this study were obtained during four intensive field campaigns in the greater Rome area, Italy (Table 1). The city of Rome covers an area of 1300 km^2 and has a population of about 2.9 million residents (about 4.3 million residents within the whole metropolitan area of 5350 km^2). Three airports are located in the Rome province, including the largest one in Italy (Fiumicino). Heavy industries are not found in Rome; the economy is mainly based on services, education, construction, tourism, etc. Parks and gardens cover some 34 % of the city area. Rome is characterised by high traffic volume and density: about 50 % of the population commutes every day, mainly by private cars. The cars per capita ratio in the city is 550 per 1000 inhabitants.

Measurements referred to as DIAPASON were performed during Oct/Nov 2013 and May/June 2014 at the Institute of Atmospheric Sciences and Climate (CNR-ISAC) in Tor Vergata, Rome. The institute is located in the south-eastern outskirts of the city (41°50’30.2’’N, 12°38’51.2’’E, 103 m a.s.l., 14 km from central Rome) and considered as an urban background site. The measurement platform MoLa (see Sect. 2.2) was positioned at a free field with no buildings within a radius of 200 m. A frequently used street is located at approximately 100 m distance in northern direction. The closest highway (A1) is situated south-westerly at a distance of about 700 m. Single-house suburbs are scattered over this territory, starting some 1 km from the site. Frascati, a town on the Alban Hills, is located at about 4 km distance in south-easterly direction. During both periods measurements at Tor Vergata were supported by the EC-LIFE+ project DIAPASON (Desert-dust Impact on Air quality through model-Predictions and Advanced Sensors ObservatioNs), which aims on improving existing tools to assess the contribution of Saharan dust to local PM_{10} levels (http://www.diapason-life.eu/, last access 11.11.2016). For this reason measurements were scheduled in periods where a dust advection event could be expected and was forecasted by a number of dust forecasts such as the DREAM8b (Basart et al., 2012), the SKIRON (Kallos et al., 1997) and the Tel Aviv University (Alpert et al., 2002) models.

The “POPE” (Particle Observations around St. PEter’s) measurement campaigns were conducted during November 2013 and June 2014 in central Rome. Measurements were performed inside a courtyard belonging to the administration of the hospital “Santo Spirito” (41°54'04.3"N, 12°27'41.5"E, 18 m a.s.l.), which is positioned approximately 600 meters from St. Peter’s Basilica. This urban measurement site is surrounded by highly frequented streets, separated from the courtyard by the four-storey building of the hospital. The surrounding area is a touristic hotspot with frequent religious gatherings (e.g. festivals, masses) and many restaurants and shops. Especially on Wednesdays during the papal audience and on Sundays, if the masses are held at St. Peter’s or during papal speeches (Angelus), the area attracts numerous visitors.

The distance between the two measurement sites is around 17 km. During both years measurements at Tor Vergata and central Rome were performed subsequently.

2.2 Instrumentation

All measurements were performed with the **Mo**bile aerosol research **La**boratory MoLa (Drewnick et al., 2012). MoLa is based on a regular Ford Transit delivery vehicle equipped with instruments for on-line measurements of chemical and physical properties of aerosols, important trace gases and meteorological variables (Table 2). Further description as well as details of the aerosol inlet system can be found in Drewnick et al. (2012). All results presented in this study were obtained in stationary measurements, with the aerosol inlet and a meteorological station at 7 m above ground level.

An HR-ToF-AMS (Aerodyne Research, Inc.; DeCarlo et al., 2006) was used to measure particulate mass concentrations of submicron non-refractory organics (“Org”), sulphate (“SO_4”), nitrate (“NO_3”), ammonium (“NH_4”) and chloride (“Chl”). The HR-ToF-AMS allows the distinction between different ions at the same nominal mass-to-charge-ratio (*m/z*). The instrument was run in V-mode, i.e. the ions followed a “V”-shaped trajectory through the mass spectrometer, allowing high sensitivity at slightly lower mass resolution, compared to the higher resolution mode (W-mode).

In the framework of the EC-LIFE+ project DIAPASON additional measurements were performed at the Tor Vergata measurement site, which aimed at assessing the contribution of Saharan dust to PM levels. These measurements included hourly PM_{10}, a three-wavelength nephelometer, one-hour filter sampling for off-line PIXE analysis (Lucarelli et al., 2014) and a polarization LIDAR-ceilometer for the assessment of presence, phase and altitude of aerosol layers (Gobbi et al., 2004). Boundary layer heights were determined from polarisation LIDAR-ceilometer measurements based on the method described by Angelini and Gobbi (2014).

Since the POPE measurements were performed inside a courtyard surrounded by four-storey tall buildings, wind speed, wind direction and solar radiation data are affected and not used for these periods. The time resolution for all measurements is 60 s or better.

3 Data preparation and analysis

3.1 General data analysis

All measured variables were corrected for sampling delays and set on a common 1-second time base. Particle losses during the transport through the inlet system were negligible (Drewnick et al., 2012). The data time series were carefully inspected and quality checked. Data affected by instrument calibrations and malfunctions were removed. Measurement periods influenced by local emissions (e.g. moving vehicles in the immediate vicinity of MoLa) were identified based on prominent short peaks in the time series of CO_2 and particle number concentration (PNC) which significantly exceeded the typical variability, and removed from the data set. After data decontamination, 5-minute averages were calculated for all variables, which were used for all following analyses if not otherwise indicated.

Data collected during the DIAPASON2013/POPE2013 and DIAPASON2014/POPE2014 field campaigns are presented in local winter (UTC+1) and local summer (UTC+2) time, respectively. For convenience, DIAPASON2013 data are presented only in winter time, even though the change from summer to winter time was at the fifth day of measurements (27.10.2013). This means data measured prior to the time change is 1 hour shifted to the past with respect to local (summer) time. Especially diurnal patterns dominated by anthropogenic activity patterns (e.g. traffic during rush hour times) could be affected by ignoring the time change. In order to evaluate this possible influence, diurnal cycles measured before and after the time change were compared, but no significant shift of diurnal patterns was observed between the two time periods. Since diurnal cycles are not only modulated by the source emission strengths, but also by boundary layer dynamics, we assume the missing evidence of the time shift in the data is caused by a combination of influence from boundary layer dynamics and the temporal uncertainty of diurnal cycles calculated over only a few days. Additionally, anthropogenic activities could have partially not been instantly adapted to the time change, which would lead to a blurring of the effect of the time change on diurnal cycles.

Polar plots of species concentration as a function of local wind direction and wind speed were generated by averaging species concentrations (60 s data) into bins of 5° wind direction and 0.5 m s^{-1} wind speed. The resulting data were smoothed by applying a natural neighbour interpolation (Sibson, 1981). As presented by Yu et al. (2004), such polar plots can provide directional information on sources in the vicinity of a monitoring site. Sources close to the measurement site are typically indicated by concentration decreases with increasing wind speed, while pollutants which are emitted from remote sources or at higher altitudes need higher wind speeds to be transported to the monitoring site (Yu et al., 2004). Similarly, Carslaw et al. (2006) reported the capability of such bivariate polar plots to distinguish between no-buoyancy sources like traffic (decreased pollutant concentration with increasing wind speed) and buoyant plumes emitted from sources like chimney stacks (increased concentrations with increasing wind speed), where the plume needs to be brought down to ground-level from a higher altitude.

3.2 HR-ToF-AMS data analysis

AMS data evaluation was performed within Igor Pro 6.37 (Wavemetrics) with the standard AMS data analysis software SQUIRREL 1.55H and PIKA 1.14H. Elemental ratios calculated from organic ion fragments (Aiken et al., 2007) based on the improved calibration method (Canagaratna et al., 2015) were determined using APES light 1.06 (all available at http://cires1.colorado.edu/jimenez-group/ToFAMSResources/ToFSoftware/). For all data sets a collection efficiency of 0.5 was applied, which is typical for the given ambient measurement conditions (Canagaratna et al., 2007). The ionisation efficiency (IE) of the ion source and the relative ionisation efficiency (RIE) for ammonium and sulphate (e.g. Canagaratna et al., 2007) were determined before the DIAPASON and after the POPE campaigns in both years. An additional IE calibration in 2013 after the field measurements showed no general trend in IE values. Therefore, the observed variability of the IE values (2013: ~10 %, 2014: ~20 %) is assumed to stem only from the uncertainty of the calibrations, and for each year averages of the determined IE and RIE values were used for data analysis. Measurements of particle free air were carried out multiple times during the campaigns and were used for correction of instrumental background effects.

In order to separate total OA into different aerosol types, PMF (Paatero and Tapper, 1994; Ulbrich et al., 2009) was applied to high-resolution mass spectra of the OA fraction with *m/z* below 131. This was done separately for each measurement campaign. The procedure of HR data and error matrices (matrix input sizes can be found in the supplement) preparation is described in detail in DeCarlo et al. (2010). Isotopes constrained to a fractional signal of their parent ion were excluded from the analysis. Within the PMF Evaluation Tool v2.06 ions with signal-to-noise ratio < 0.2 were removed from data and error matrices, and ions with signal-to-noise ratio between 0.2 and 2 were down-weighted in the analysis by increasing their estimated error by a factor of two (Ulbrich et al., 2009). Particulate CO_2^+ (*m/z* 44) and its associated ions at *m/z* 16, 17, 18 and 28 were down-weighted by a factor of $\sqrt{5}$ (Ulbrich et al., 2009, supplemental information).

In order to find the most reasonable and robust PMF solution, the number of factors (one up to ten, always at least two more than the finally selected solution), the rotational force parameter (fPeak: -1 to 1; $\Delta = 0.2$) and the starting point (seed: 0 to 50; $\Delta = 1$) were varied (see Ulbrich et al., 2009 for methodological details). Solutions with fPeak=0 and seed=0 turned out to yield robust results for all data sets. The evaluation of potential PMF solutions was based on comparisons of the resulting factor time series with those of co-located measurements (see Sect. 4.2), and of factor mass spectra with such from the literature. Residues, i.e. the contribution of organic mass concentrations not included in any of the factors, accounted for <1 % of the organics mass concentration in all used PMF solutions and are therefore negligible. Mass spectra and time series of each identified PMF factor can be found in the supplement.

4 Results and discussion

4.1 Overview: Differences between seasons and locations

This section provides a broad overview of the mean conditions of local meteorology and air quality during each measurement campaign (Table 3, Fig. 1), and discusses their seasonal and spatial differences. Figure 2 provides an overview of the relative composition of non-refractory PM_1 plus BC and shows the contribution of different factors related to different sources retrieved from the OA fraction using PMF. In total, seven different factors were identified: OOA (oxygenated OA), SV-OOA (semi-volatile OOA), LV-OOA (low-volatile OOA), HOA (hydrocarbon-like OA), COA (cooking OA), BBOA (biomass burning OA) and CSOA (cigarette smoke OA, see Sect. 4.2.5; considered local contamination and not included in the pie charts in Fig. 2). Here, only a general overview of these different factors focusing on seasonal and spatial differences is given. A more detailed discussion of the various sources associated with these factors is provided in Sect. 4.2.

Meteorology overview:

The first period of the DIAPASON2013 campaign (23.-31.10.2013) was dominated by high pressure influences and low wind speeds with air masses moving from the Atlantic across Northern Africa and the Mediterranean basin. Within this period dust from the Sahara was transported to the Rome area (see Sect. 4.2.1). The second half of DIAPASON2013 (1.-7.11.2013) was characterized by a pressure drop and increased wind speed together with some frontal passages leading to precipitation. Both turbulent kinetic energy (TKE), which is a measure of the intensity of turbulence and can be used as an indicator for the mixing efficiency of pollutants in the air (Srivastava and Sarthi, 2002), and boundary layer heights (BLH) were at rather low levels during DIAPASON2013 (Table 3), favouring the accumulation of pollutants.

The first measurement days of POPE2013 (07.-09.11.2013) were influenced by changing weather conditions, followed by a low pressure system centred over Italy (10.-14.11.2013) driving N-NW wind conditions and leading to almost daily precipitation events. High TKE levels (almost three times higher compared to DIAPASON2013) and slightly increased boundary layer heights (900 m compared to 850 m; Table 3) led to conditions were dilution of pollutants was more favoured.

During DIAPASON2014 a low pressure system was located over the Atlantic and North Africa. Saharan dust was advected to the measurement site during the first week of measurements (Rizza et al., 2016). Some precipitation events occurred during these advections. TKE was slightly increased after the dust advection, favouring the reduction of pollutant concentrations. Boundary layer heights reached around 1500 m (Table 3).

During POPE2014 a period of low pressure over the Atlantic and high pressure over Africa and Europe (04.-13.6.2014) was followed by reversed conditions (14.-17.6.2014) with some heavy precipitation events. Compared to DIAPASON2014 the boundary layer was slightly higher (1560 m) and TKE was decreased by 20 %.

During DIAPASON2013 local winds were predominantly arriving from south and south-easterly directions, whereas during DIAPASON2014 south-westerly wind directions dominated (Fig. 1; no data available for POPE2013/2014, see Sect. 2.2). No clear relationships of air mass origin with measured PM_1 mass concentrations were found by the calculation of HYSPLIT (Stein et al., 2015) and FLEXPART (Stohl et al., 2005) backward trajectories for our measurement periods.

Aerosols and trace gases:

During both Oct/Nov and May/June measurements, CO_2, NO_x and particulate PAH concentrations (all typically traffic-related) were higher in central Rome compared to the outskirt location (Table 3). Cooking-related aerosol (COA) concentrations were found to be generally higher in the city centre (Fig. 2b, d), while traffic-related (HOA) concentrations were strongly increased during 2014 (+ 58 %) at the urban location (Fig. 2d), and nearly the same at both locations during the 2013 measurements (Fig. 2a, b). All this is consistent with increased primary emissions (cooking, traffic) at the urban (Fig. 2b, d) compared to the suburban site (Fig. 2a, c). However, total concentrations of organic aerosol predominantly of primary origin (POA) were higher at the suburban (1.9-2.5 $\mu g\ m^{-3}$) compared to the urban (1.4-1.6 $\mu g\ m^{-3}$) location. This is due to a factor indicating particles from biomass burning (BBOA), which was obtained exclusively at Tor Vergata (Fig. 2a, c). Here, biomass burning seems to be an important particle source, contributing the most abundant fraction (42-51 %) of POA. Note that due to the fact that OOA concentrations are significantly higher in the suburban site in 2013, the relative contribution of POA-related aerosol types (i.e. HOA, COA, and BBOA) is lower at the suburban location during this year. However, in order to avoid a bias in the comparison of the more locally generated POA-related aerosol types by advected aerosol mass, we compare the absolute mass concentrations for the different aerosol types at the different measurement locations/times. BC, which is related to primary emissions from both biomass burning and traffic, showed no general trend between the two locations (Table 3). Ratios of HOA/BC (DIAPASON2013: 0.26, POPE2013: 0.33, DIAPASON2014: 0.46, POPE2014: 0.53) were lower during fall 2013, indicating the more dominant contribution of biomass burning emissions to BC during this time. The fact that contributions from biomass burning to total BC concentrations are non-negligible was also found earlier (e.g. Crippa et al., 2013a). Generally, the HOA/BC ratios were higher in the urban compared to the suburban location in each year. This reflects the stronger influence of traffic emissions and the lower contribution of biomass burning emissions in the city centre.

The influence of increased temperatures and stronger solar radiation during measurements in May/June 2014 (Table 3) is reflected in elevated ozone mixing ratios and the fact that it was possible to extract an additional OOA factor, which was attributed to a fresher, less oxidised aerosol type (SV-OOA, Fig. 2c, d). This SV-OOA likely is the result of quick formation of secondary aerosol from precursors originating from the greater Rome area. Within a continuum of OOA with different degrees of aging/oxidation, SV-OOA (fresh, of rather local origin) and LV-OOA (aged, of more regional origin) are located in the upper and lower range, respectively. In contrast, during Oct/Nov, only one type of rather aged OOA was found (Fig. 2a, b), due to reduced photochemistry in this season which prevents the quick formation of oxygenated aerosol from precursors emitted in the vicinity.

A stronger influence of aged aerosol of rather regional origin on the PM_1 fraction was observed for May/June than for Oct/Nov: the fraction of OOA (SV-OOA + LV-OOA) to OA, but also the relative contribution of total organics and sulphate to PM_1 were higher in the warmer season (Fig. 2c, d). This could be caused by different prevalent air mass origins, though the analysis of backward trajectories provided no definite answers on this (see above). Also the relative fraction of sulphate could be lower in Oct/Nov due to an enhanced contribution of primary particles (Table 3) as a consequence of lower boundary layer heights (which limits the dilution of locally produced aerosols while it does not influence the concentration of transported aerosols) and, potentially, higher emission strength of local sources during the colder season. Consistently, in Oct/Nov 2013 at both measurement locations a higher BC fraction was observed compared to May/June, and POA made up a larger fraction of the total measured organics (Fig. 2). Also absolute BC concentrations were enhanced. Additionally, higher concentrations of NO_x and PAH (increased by a factor of 3), and of total particle number concentration (PNC) were observed during Oct/Nov (Table 3), indicating the accumulation of pollutants during the colder season due to the aforementioned reasons.

"Home-made" vs. "advected" PM_1:

For a rough estimate of the contribution of PM_1 originating from sources in the Rome metropolitan area and from advection from outside, PM_1 species were separated into "home-made" (BC, HOA, COA, BBOA, SV-OOA, NO_3) and "advected" (OOA/LV-OOA, SO_4)

(Table 4). NH_4 was apportioned to home-made and advected PM_1 based on the molar concentrations associated with NO_3 and SO_4, respectively. Not included in these estimates are dust advection periods and emissions from cigarettes (which were considered as local contaminations).While nitrate is formed quickly and thus strongly contributes to the "home-made" aerosol fraction, it can also be transported. The resulting fraction of "home-made" aerosol might therefore be slightly biased high, however by a few percent at most due to the small total contribution of nitrate to PM_1. For the POPE2013 measurements OOA shows only poor correlations with SO_4, but slightly better ones with NO_3, suggesting either a local contribution of OOA or the transport of NO_3. Whatever the case may be, it will contribute to the error of the estimation. During the measurements the fraction of home-made PM_1 accounted for 42-70 % and advected PM_1 accounted for 30-58 % of total (home-made + advected) PM_1 (Table 4). During the 2013 measurements higher absolute concentrations of home-made PM_1 were found at the suburban location compared to the urban site, possibly caused by meteorological conditions favouring pollutant accumulation during the respective period (see above). However, due to the large differences in total PM_1, the relative contribution of home-made PM_1 was higher at the urban location. During 2014, when meteorological conditions (e.g. BLH, TKE; see Table 3) were rather comparable at both locations, similar absolute concentrations of home-made PM_1 were observed at both sites. No general seasonal difference in home-made PM_1 fraction was observed, although BLH was strongly increased during the May/June 2014 compared to the Oct/Nov 2013 measurement periods. Partially, this might be due to the additional presence of the home-made species SV-OOA during May/June, which could to some extent have compensated for the dilution effect due to the increased BLH. Altogether, neither a general spatial (DIAPASON vs. POPE) nor a seasonal (Oct/Nov vs. May/June) tendency regarding the contribution of home-made and advected PM_1 to total PM_1 was observed. These results indicate that urban air quality in Rome strongly depends on both, emissions within the city and transport of pollutants to the city, which both contribute to urban aerosol concentration.

Independent of measurement season and location the organics fraction was always found to contribute the largest share of PM_1 (44-53 % of non-refractory PM_1 plus BC, Fig. 2), though its relative composition (primary vs. secondary OA) differed. Regarding absolute PM_1 concentrations (Table 3 from EDM measurements, Fig. 2 from non-refractory components plus BC) neither any general conclusion whether aerosol mass concentrations are higher at the city centre or in the suburb, nor whether PM_1 concentrations are elevated during any of the two different seasons, can be drawn. In the 2013 campaign total PM_1 mass concentrations were more than doubled at the suburban compared to the urban location, whereas in the 2014 measurement PM_1 concentrations were increased by a factor of 1.4 at central Rome. As discussed above, changes in meteorological conditions are likely one explanation for this result: During DIAPASON2013 meteorological conditions favoured the accumulation of pollutants, whereas the dilution of pollutants was favoured during the POPE2013 measurement period; during DIAPASON2014 and POPE2014 TKE and BLH were rather similar leading to comparable pollutant dilution effects during the two measurement periods. BLH were increased by around 75 % during May/June compared to Oct/Nov 2013, leading to stronger dilution capacities in general. In addition to meteorological conditions (e.g. solar radiation, BLH, TKE, air mass origin, etc.) local air quality can be strongly influenced by local emission from various sources (traffic, cooking, biomass burning). A strong influence of meteorological conditions (air mass origin) on air quality was also observed during the MEGAPOLI campaign in Paris in July 2009, where variations in secondary aerosol concentration mainly were attributed to such reasons (Freutel et al., 2013).

4.2 Aerosol sources: identification and characterisation

In this section the various aerosol types and sources which were identified from the data obtained during the DIAPASON and POPE measurement campaigns in 2013 and 2014 are discussed in more detail. Each identified aerosol type was characterized in an attempt to determine its contribution to total particulate mass and its seasonal and spatial variability. Furthermore, the potential origin of the identified aerosol types is discussed.

4.2.1 Saharan dust

During each DIAPASON field campaign one dust advection event lasting for several days was observed. The identification of the dust events with dust reaching down to the ground was based on dust forecasts provided by the SKIRON model (Kallos et al., 1997) and on co-located polarisation LIDAR-ceilometer measurements (data can be found at www.diapason-life.eu). PIXE analysis of 1-hour filter samples confirmed a significant increase of mineral dust concentrations (i.e. Na, Mg, Si, Al, Ti, K, Ca, Fe; Nava et al., 2012) at ground level during the identified dust periods (Barnaba et al., in preparation 2016). Table 5 provides the time intervals of “dust” and “no dust” (i.e. background with respect to dust advections) periods for DIAPASON2013 and DIAPASON2014. Based on these, estimations regarding the contribution of dust to PM levels were made by calculating the coarse particle fraction ($PM_{10-2.5}$) from EDM measurements for the respective periods (Table 5).

During DIAPASON2013 a Saharan dust advection period was observed from the beginning of the measurements (October 23th) until November 1st with $PM_{10-2.5}$ concentrations at ground level being increased by 80 % with respect to background conditions. Total PM_{10} concentrations were elevated by 150 %, with 71 % and 57 % of mass belonging to PM_1 during the dust event and during the background periods, respectively. This increase in absolute PM_1 with an increase in the fraction of PM_1 during the dust advection compared to background conditions was also reflected in the particle size distributions (Fig. 3, black traces). The dust event was also forecasted by the BSC-DREAM8b model (Basart et al., 2012; Perez et al., 2006a; Perez et al., 2006b) and also HYSPLIT (Stein et al., 2015) back trajectories showed dust transport from the Sahara with main dust sources located at 30-33N, between Morocco (Saharan Atlas) and SW Tunisia (Erg Oriental), in the period 26.-28.10.2013, moving along an anti-cyclonic pattern.

During DIAPASON2014 a dust event was observed from May 20th (1 pm) until May 26th (9 am). Here, the coarse particle fraction ($PM_{10-2.5}$) was increased by 180 % compared to the "no dust" interval with much smaller fractions of particle mass in PM_1 (31 and 48 % in the "dust" and "no dust" periods, respectively). The contribution of dust to the coarse particle fraction with almost no contribution to the submicron fraction is reflected in the difference particle mass size distribution (Fig. 3, bottom panel, red trace), which shows maximum mass concentrations for aerodynamic particle diameters around 2 µm and smaller contributions extending down to ~600 nm and up to more than 10 µm particle diameter. Comparing both dust events in terms of particle sizes, the contribution of the dust advection event during DIAPASON2014 was characterized by a broad particle mass size distribution with maximum concentrations at particle sizes around 2 µm, whereas in 2013 two modes (with maxima around 0.6 µm and 3 µm) were observed. Both the BSC-DREAM8b model and HYSPLIT back trajectories showed dust transport from the Sahara to occur between 19.-22.05.2014 along a cyclonic pattern, with dust originating (as in October 2013) at 30-35N between Morocco (Saharan Atlas) and SW Tunisia (Erg Oriental).

With regard to the coarse particle mode ($PM_{10-2.5}$) the dust event during DIAPASON2014 was more distinct. In terms of absolute PM_{10} concentrations, higher concentrations at ground level were reached during the dust advection measured during DIAPASON2013. However, with respect to the "no dust" conditions, mean PM_{10} increases of about 17 and 11 µg m^{-3} were observed during DIAPASON2013 and DIAPASON2014, respectively. This agrees well with the observations made by Gobbi et al. (2013), who reported PM_{10} concentrations in the order of 12-16 µg m^{-3} during dust advection events in the period 2001-2004. During both advection events legal PM_{10} limits of the European Union (daily mean value of 50 µg m^{-3}) were not exceeded.

AMS data were investigated for potential impacts of Saharan dust advections on the chemical composition of non-refractory submicron particles. Figure 4 presents the mean chemical composition of non-refractory PM_1 measured during “dust” and “no dust” periods for both DIAPASON campaigns. Only in 2013 differences in absolute mass concentrations were observed, consistent with higher submicron particle concentrations observed in the mass size distributions during this dust event (Fig. 3, bottom panel, black trace). This difference in PM_1 concentration between the dust and no dust period is likely due to differences in air mass origin. Whereas during the dust period air masses preferentially arrive from the Mediterranean region, the no dust period is mainly influenced by air masses from the Atlantic Ocean. However, the relative composition remains nearly the same for both periods in both years. This result suggests that there was no significant influence of the dust advection on the chemical composition of the submicron non-refractory aerosol fraction.

In order to cross-check whether differences in meteorological conditions could have biased these results, and e.g. could have compensated for changes due to the dust affecting non-refractory PM_1, averages of meteorological variables were calculated for “dust” and “no dust” periods. The only slight differences found between dust events and background conditions were in local wind directions (SE compared to S in 2013, and SW compared to SSW in 2014) and (in 2013) in wind speeds ((1.5 ± 0.8) m s^{-1} compared to (2.7 ± 1.6) m s^{-1}). Since these differences are only very minor, we conclude that our observation of comparable chemical composition of non-refractory PM_1 during “dust” and “no dust” periods was not caused by any compensating effects.

4.2.2 Secondary and aged aerosol: seasonal influence on formation processes and chemical composition

Seasonal variations of the characteristics of secondary and aged aerosols were identified by investigating new particle formation events and particle chemical composition measured during both POPE and DIAPASON campaigns in Oct/Nov 2013 and May/June 2014.

New particle formation:

Diurnal cycles of size-resolved and total particle number concentrations (PNC) for the Oct/Nov 2013 and May/June 2014 campaigns (Fig. 5) revealed several seasonal differences. During measurements in Oct/Nov 2013 (Fig. 5, left panels), elevated PNC were only observed during rush hour times in the morning and the evening, whereas in the measurements in May/June 2014, an additional PNC peak occurred with a maximum around 1-3 pm (DIAPASON2014) and 2-3 pm (POPE2014), respectively (Fig. 5, right panels). At the urban site (Fig. 5, bottom panels), PNC after the morning rush hour remain at an elevated level, likely because of increased background concentrations due to generally higher traffic density in the city centre. PNC were generally higher in the Oct/Nov 2013 than in the May/June 2014 campaigns, as discussed in Sect. 4.1.

Mean particle number size distributions for the period of maximum PNC at midday (Fig. 6; "nucleation"; solid traces) show a distinct mode at small particle diameters between 7-15 nm for both May/June 2014 campaigns (also visible in Fig. 5, right panels) compared to the number size distribution measured between 10-11 am (Fig. 6; "background"; dashed traces). Such occurrences of ultrafine particles at midday, when concentrations of particles from traffic are at a relative minimum and thus not responsible for strongly increased PNC, have been attributed before to new particle formation characteristic for urban areas with high solar radiation (Brines et al., 2015; Minguillon et al., 2015; Reche et al., 2011). During POPE2014 an additional mode at larger particle sizes (D_p approximately 15-50 nm) was found in the particle size distribution measured at midday (Fig. 6), probably originating from increased background levels.

During POPE2014 diurnal cycles of mean (grey) and median (black) PNC agree very well with each other, also during the midday peak (Fig. 5). This reflects the observed low day-to-day variability during this period for the measurements in central Rome, wherein the corresponding PNC time series a midday peak was observed on every single day. Local smoking activities (see Sect. 4.2.5) at the central Rome measurement location seem not to have biased these results, since no differences in diurnal cycles of PNC for weekdays (smoking activities) and weekends (no smoking activities) have been found. This suggests that the formation of new particles around midday was taking place every day at central Rome during the May/June 2014 measurement period.

Conversely, the diurnal cycle of total PNC measured during DIAPASON2014 shows a clear discrepancy between mean and median values during the midday peak (Fig. 5). This discrepancy is due to the fact that new particle formation events did not occur on all days, probably induced by different meteorological conditions and/or differences in pre-existing particle surface areas (e.g. Kulmala and Kerminen, 2008).

To test whether particular meteorological conditions can promote/suppress new particle formation events, “nucleation” and “non-nucleation” days were classified for DIAPASON2014 by comparing PNC measured during 10-11 am (background conditions, PNC_{bg}) and during 11 am - 4 pm (typical nucleation periods, PNC_{nuc}) (Table 6). This classification was cross-checked by verifying if during classified nucleation days a clear increase in PNC at small particle diameters ($D_p \leq 25$ nm) could be observed in the particle number size distributions, and whether it was missing on classified non-nucleation days. Only one potentially falsely classified nucleation day (24.05.2014) was found by checking these criteria, and was moved to the class of non-defined days. The classification resulted in six nucleation days, six non-defined days and two non-nucleation days.

Mean values for the time period 10 am to 4 pm (new particle formation period plus one previous hour) were calculated for each day and averaged according to the above-mentioned classification for the DIAPASON2014 campaign. Table 6 lists PNC of the classified periods and variables potentially supporting new particle formation. A slight trend of increased temperature, solar radiation and ozone levels and of low relative humidity characterizes nucleation days compared to non-nucleation days and non-defined days. This is consistent with previously reported association of high solar radiation (Pikridas et al., 2015; Shi et al., 2001), low relative humidity (Kulmala and Kerminen, 2008) and increased ozone concentrations (Harrison et al., 2000) with new particle formation events. No relationship between the occurrence of new particle formation and the presence of Saharan dust was observed. Different to the findings of Zhang et al. (2004), no increase of sulphate, ammonium and nitrate concentrations was observed in our measurements during periods with new particle formation events. Estimations based on the size distribution measurements during DIAPASON2014 reveal that less than 1 % of PM_1 can be assigned to particles generated by new particle formation. Therefore, it is not surprising that no significant influence of the particle formation events on the AMS-measured chemical particle composition could be observed.

In summary, our data do not provide sufficient statistical evidence to unequivocally determine the driving factors for new particle formation. However, since indications for such were only observed in the warmer season at both locations, it is probably linked to higher temperatures and stronger solar radiation. Similar results were obtained from the MEGAPOLI measurements in Paris, where new particle formation was only observed during summer (Pikridas et al., 2015). During the May/June 2014 campaigns, new particle formation events occurred roughly on 43 % of the measurement days at the suburban location, but on each single day at central Rome, potentially due to increased concentrations of precursors and higher prevailing mean temperatures (Table 3). In a long-term measurement study performed by Costabile et al. (2010) the occurrence of aged nucleation mode particles (up to 30 nm) was observed predominantly in spring in the area of Rome at a regional background site (located more remotely than the Tor Vergata site) in the early afternoon (3 pm) when the measurement site was located downwind of Rome (Brines et al., 2015). In contrast, during DIAPASON2014 measurements, no dependency between nucleation events and wind direction was observed, and the site was not located downwind of Rome during nucleation periods. At this measurement location, probably lower concentrations of precursors are available than in central Rome, but higher concentrations than at a remote regional background site such as in the study of Costabile et al. (2010). This probably facilitates new particle formation events in the direct vicinity of the site under favourable meteorological conditions, but independent of advection of air masses from central Rome.

Secondary and aged aerosol:

Because of extremely low mass contributions from freshly formed particles to total particle mass, new particle formation had no influence on the measured total organics mass concentrations. However, a general seasonal difference in the composition of the oxygenated organic aerosol (OOA) as determined in the PMF analysis was found, as discussed in the following.

OOA, an aerosol type with increased oxygenation level, typically dominates the OA fraction. It is assumed to be mainly formed in the atmosphere from gaseous biogenic and anthropogenic precursors by photochemical oxidation, thus indicating SOA. Additionally, some OOA may originate from atmospheric aging of POA. Generally, aging processes are reflected in an increased degree of aerosol oxidation (Jimenez et al., 2009) leading to a larger fraction of m/z 44 (CO_2^+) in the aerosol mass spectra, generated by thermal decomposition of carboxylic acids in the AMS (Alfarra et al., 2004). Additionally, a prominent peak in OOA mass spectra occurs at *m/z* 43 ($C_3H_7^+$ and $C_2H_3O^+$). Under conditions where sufficient freshly oxidised organic aerosol is available in the ambient air, PMF can separate the OOA into two factors. These factors differ in their relative fractions of *m/z* 44 and *m/z* 43 (f44 and f43, ratio of *m/z* 44 and *m/z* 43 signal, respectively, to the total signal of organics), which reflects their different degrees of oxidation. The factor associated with higher f43 indicates a less oxidised, fresher, more locally produced semi-volatile OA (SV-OOA), whereas a higher f44 indicates a stronger oxidised, more aged low-volatile OA (LV-OOA) (Ng et al., 2010).

During our measurements the relative contribution of OOA to total organics varied between 42-76 %, with a slight increase during the warm period (Fig. 2). However, the main seasonal difference was found in the composition of the OOA fraction. During the Oct/Nov 2013 campaign only one type of OOA was found, whereas for May/June 2014 PMF analysis resulted in two OOA-factors: SV-OOA and LV-OOA. This is typically only observed during summer conditions, when the dynamic range of temperature, ozone concentration and solar radiation is large, which is assumed to be the main driving force for the variability of the OOA volatilities (Jimenez et al., 2009). Similar observations were made during the MEGAPOLI measurements in Paris, where only one factor describing OOA was identified in winter (Crippa et al., 2013a), whereas during summer SV-OOA and LV-OOA could be separated (Crippa et al., 2013b).

The ratio of f44 to f43 gives an indication on the mean aging level of the aerosol (Fig. 7; Ng et al., 2010). LV-OOA (from DIAPASON and POPE 2014, red markers) and OOA (from DIAPASON and POPE 2013, green markers) fall into the same region in the f44 vs. f43 plot (Fig. 7), indicating similar aging stages. SV-OOA (from DIAPASON and POPE 2014, blue markers) shows a much higher fraction of *m/z* 43 together with a decreased *m/z* 44 fraction, which suggests a low-oxidised, less aged particle type. Also the recombined “LV-OOA+SV-OOA” (black markers in Fig. 7) shows a stronger contribution of f43 compared to OOA, indicating an overall higher fraction of less oxidised organic aerosol in the warmer season. We assume that LV-OOA (and OOA) is mainly advected and consists of strongly processed material, whereas the low oxidation level of SV-OOA suggests a fresh, more locally produced aerosol which was quickly formed from regional precursors as a consequence of increased photochemistry during this season.

This hypothesis is tested by the use of polar plots, which connect species concentration with information on local wind direction and speed, thereby indicating the origin of a certain type of aerosol (see Sect. 3.1). Figure 8a shows the colour coded concentration of SV-OOA, LV-

OOA, NH_4, NO_3 and SO_4 depending on wind direction and speed obtained during the DIAPASON2014 measurements. SV-OOA concentrations are increased during low wind speed conditions, indicating nearby, no-buoyance sources. In contrast, LV-OOA concentrations are almost independent of wind speed with only slightly increased concentrations during periods of high wind speed with mainly south-westerly wind directions (direction of Tyrrhenian Sea). This suggests that LV-OOA is not associated with sources located in the vicinity of the measurement site, but long-range transported to the site e.g. over the ocean or from central Europe. The polar plot characteristics of NO_3, which is often used as a tracer for semi-volatile aerosol (DeCarlo et al., 2010; Lanz et al., 2007), show strong similarities to the ones of SV-OOA (Fig. 8a). Also SO_4 and LV-OOA show polar plot patterns similar to each other (Fig. 8a), confirming the characteristics of an aged, regionally transported aerosol. The polar plot of NH_4 shows a hot spot at low wind speeds in northern direction, which is also reflected in the patterns of SV-OOA and NO_3. Increased NH_4 concentrations are also observed at higher wind speeds in south-westerly direction, agreeing with the polar plot patterns of LV-OOA and SO_4.

The polar plot of OOA obtained for DIAPASON2013 shows increased concentrations particularly during periods of north-easterly, but also during south-westerly wind directions (Fig. 8b). During conditions of low wind speed, OOA concentrations are increased independent of the prevailing wind direction. In contrast to the findings for DIAPASON2014, for this data set similar polar plot characteristics as for OOA were observed partly in the plots of NO_3, NH_4 and SO_4 (Fig. 8b). Elevated NH_4 and SO_4 concentrations were mainly measured during times with south-westerly wind direction, whereas for NO_3 rather an increase for north-easterly directions was observed. Based on the polar plot characteristics no consistent trend indicating the degree of aging, the source or the formation process of the OOA fraction can be observed, consistent with the assumption of advection of a rather aged type of OOA together with different amounts of NO_3, SO_4 and NH_4 depending on air mass history.

4.2.3 Particles from biomass burning

The type of primary organic aerosol at Tor Vergata identified from PMF analysis which had the largest share during both measurement periods was attributed to biomass burning (biomass burning OA, BBOA). BBOA was identified by comparison with the time series of known ion fragments of levoglucosan ($C_3H_5O_2^+$ (*m/z* 73) and $C_2H_4O_2^+$ (*m/z* 60); Schneider et al., 2006) and by correlating the BBOA mass spectra with those presented by Mohr et al.

(2012). With Pearson's R^2 = 0.57-0.59 rather poor correlations were obtained, which demonstrates the complexity and the potential variations of the BBOA mass spectra due to aging processes and differences in source processes (e.g. different burning conditions or fuels; Weimer et al., 2008). The mean BBOA mass concentration was 1.28 µg m^{-3} (24 % of total OA) during Oct/Nov 2013 and 0.82 µg m^{-3} (18 % of total OA) during May/June 2014, respectively (Fig. 2). Increased BBOA concentrations during the colder season probably result from stronger agricultural burning activities (green waste burning) and potentially domestic heating, as well as from lower boundary layer heights. However, due to the moderate temperatures also during the Oct/Nov 2013 measurements, we do not expect strong contributions from domestic heating. In the evening of 25.10.2013 during a strong, visually detectable biomass burning event, which could be related to green waste burning in the nearby Alban Hills, maximum concentrations of BBOA were obtained (up to 75 µg m^{-3}). This event was used during the identification of the PMF solution, since only a factor including this event could be considered to be attributed to biomass burning emissions.

The origin of BBOA emissions was further investigated by relating BBOA mass concentrations to local wind direction and speed (see Sect. 3.1). The resulting polar plots (Fig. 9) indicate BBOA particles mainly arriving from south-easterly directions during DIAPASON2013. During this measurement period agricultural fires were frequently observed in the Alban Hills (Frascati vineyard area), which are located in this direction. Additionally, BBOA was observed during conditions of north-easterly winds and higher wind speeds (up to 4 m s^{-1}), possibly resulting from residential wood burning in a densely populated urban periphery area (Borghesiana). The polar plot of BBOA obtained from DIAPASON2014 measurements does hardly point to any preferential direction of BBOA origin. Since elevated BBOA concentrations were mainly reached during low wind speed conditions, emission from rather local sources is suggested.

BBOA was not identified in the measurements in the city centre of Rome, even not in PMF solutions with a large number of factors (8). Since green-waste burning and domestic heating with biomass are forbidden in central Rome, biomass burning related particles are probably not emitted in the local environment. However, the contribution of biomass burning and domestic heating to the urban air pollution of Rome (especially during winter time) was reported by Gariazzo et al. (2016). Apparently, during our measurements in late spring and autumn the contribution of particles emitted outside the suburban area is too small to be identified with PMF and/or aging processes during the transport of the particles lead to a loss of the BBOA fingerprint (Bougiatioti et al., 2014). Also during the MEGAPOLI measurements in Paris emissions from biomass burning were identified in the organic aerosol fraction (Crippa et al., 2013a). In contrast to our measurements, in Paris BBOA was only found during the winter time campaign (Jan/Feb), but also at the measurement location in the city centre, probably generated by local domestic wood burning (Crippa et al., 2013a).

In summary, the results from both DIAPASON measurements show that particles from biomass burning significantly (18–24 % of total OA) contributed to local air pollution in the suburban area in late autumn as well as in late spring. Agricultural fires and possibly wild fires probably are their most important sources, since heating activities are assumed to be quite low at these times due to the moderate temperature conditions (Table 3).

4.2.4 Emissions related to traffic and cooking activities

At both measurement locations, during both seasons particles from traffic- as well as from cooking-related emissions were detected.

Indications of traffic-related emissions can be found in several measured variables showing a distinct diurnal pattern with peaks during the morning and evening rush hours. The time series and diurnal pattern of HOA (hydrocarbon-like organic aerosol), a PMF factor that is typically associated with traffic emissions, show good agreement with the respective patterns of species like BC, NO_x and PAH (diurnal cycles: $R^2 > 0.85$) for all four campaigns. Also correlations of complete campaign time series of HOA with BC result in good agreements ($R^2 \approx 0.7$).

In the diurnal cycles of HOA seasonal and spatial differences can be observed (Fig. 10). Independent of season and measurement location a short peak occurs during the morning rush hour and a broader peak starting during the evening rush hour. During all field campaigns except DIAPASON2013, HOA concentrations remain increased throughout the night. Thus, the exact period of the evening rush hour cannot be clearly isolated. These differences in the shapes of the HOA peaks in the morning and evening rush hour are mainly controlled by boundary layer dynamics together with the diurnal cycle of traffic-related emissions (rush hour times). A seasonal difference is observed in the HOA evening rush hour peak, which peaks around midnight during May/June, but around 7-8 pm during Oct/Nov. This shift and the broadening of the HOA peak in May/June 2014 is probably driven by the different boundary layer dynamics during the two seasons.

For both measurement years a time shift of the morning peak between Tor Vergata and central Rome (later by about one hour) can be observed. Since similar diurnal temperature profiles measured at the suburban and the urban location suggest also similar boundary layer dynamics at the two sites, the observed shift possibly results because traffic starts in the suburbs earlier in the morning and continues slowly towards the city centre. In contrast to our observations, from BC measurements during the MEGAPOLI summer campaign in Paris no distinct shift of the morning rush hour peak was observed between the two suburban and the urban measurement locations (Freutel et al., 2013).

Mean HOA mass concentrations for the individual measurement campaigns range between 0.59-0.93 µg m^{-3}. During the 2013 measurements (Oct/Nov) similar concentrations were obtained at the suburban site (0.76 ± 1.04 µg m^{-3}) and central Rome (0.71 ± 0.72 µg m^{-3}), whereas in 2014 higher concentrations were reached at central Rome (0.93 ± 0.73 µg m^{-3}) compared to the suburb (0.59 ± 0.60 µg m^{-3}). Overall, the contribution of traffic-related emissions (e.g. HOA, NO_x, PAH) to local air pollutant levels was higher in central Rome, as already discussed in Sect. 4.1.

A factor associated with cooking emissions, COA (cooking OA), was obtained by PMF analysis of the OA measured at both locations and during both seasons. The COA mass spectra show prominent peaks at *m/z* 41 and 55 (Allan et al., 2010; Lanz et al., 2007) and a smaller contribution of *m/z* 60 and 73 (Mohr et al., 2009). Our COA mass spectra correlated well with those found by Faber et al. (2013) and Mohr et al. (2012) with $R^2 = 0.63$-0.93.

The COA diurnal cycles observed at central Rome (Fig. 11, upper panel) are consistent with results from previous studies (e.g. Allan et al., 2010; Mohr et al., 2012) showing highest concentrations in the late evening (around 10 pm) and a smaller peak around midday (2-3 pm). This pattern is generated by a combination of source strengths and boundary layer dynamics, with typically increased boundary layer height during lunch time compared to dinner time.

In contrast, diurnal cycles of the COA factors measured at the suburban location in 2013 and 2014 (Fig. 11, lower panel) both show a peak in the evening, but only during DIAPASON2014 a slight and barely significant COA concentration increase was observed during lunch time. This could be due to an insufficient separation of the COA and HOA factor during PMF analysis, which is also demonstrated in the COA “morning peak” of the DIAPASON2014 measurements. However, the missing midday peak also reflects the generally low abundance of cooking-related OA at the suburban measurement location: while there are strong cooking activities and a large abundance and closeness of restaurants around the central Rome site, potential sources in the immediate vicinity of the suburban site are scarce. At a distance of around 250 m from our monitoring site, a cafeteria served hot meals for lunch, but apparently, our measurements were not strongly affected by its emissions.

Consistently, absolute mass concentrations of cooking-related emissions were higher at the central Rome site (0.70 ± 1.00 µg m^{-3}, 0.65 ± 0.69 µg m^{-3} in 2013 and 2014, respectively) compared to the suburban measurement location (0.45 ± 0.50 µg m^{-3}, 0.53 ± 1.29 µg m^{-3}). Ranging between 8-29 % of the total OA concentrations, cooking activities contribute significantly to (sub-) urban air pollution. During meal times the contribution of COA to total organics can be very high. For example during lunch/dinner times at central Rome, COA contribution to total organics was 35 %/53 % (POPE2013) and 9 %/25 % (POPE2014), respectively. Similar observations were made during the MEGAPOLI winter measurements in Paris, where COA contributed on average 11-17 % to total OA (up to 35 % during lunch times) (Crippa et al., 2013a).

4.2.5 Cigarette smoking emissions

For both POPE campaigns in central Rome PMF analysis of the organic aerosol fraction resulted in a factor which could be associated with cigarette smoke (CSOA; excluded from Fig. 2). This was not very surprising, since cigarette smoking took place in the direct vicinity of the measurement location. The mass spectra of CSOA from both years show good correlation with each other (R^2 = 0.7; Fig. 12). Very characteristic for the CSOA spectra is a peak at *m/z* 84 from $C_5H_{10}N^+$ and a peak at *m/z* 42 resulting from the ion $C_2H_4N^+$ (Fig. 12). Both ions are typically observed in EI mass spectra of nicotine (NIST: http://webbook.nist.gov, last access 11.11.2016). $C_5H_{10}N^+$ (*N*-methylpyrrolidine), which shows the strongest signal in the EI mass spectrum of nicotine, is generated by cleavage of the nicotine molecule into two heterocycles (Jacob III and Byrd, 1999). Since nicotine is one of the most abundant particulate compounds identified in cigarette smoke samples (Rogge et al., 1994), its fragments are suitable tracers for cigarette emissions. While cigarette smoke-related aerosol has been found in AMS measurements previously (Faber et al., 2013; Fröhlich et al., 2015) and also the detection of nicotine from cigarette smoke was mentioned (Jayne et al., 2000), to our knowledge, the identification of the nicotine fragment *N*-methylpyrrolidine from analysis of HR-ToF-AMS data is reported here for the first time. The time series of $C_5H_{10}N^+$

was used during the evaluation of the PMF results as tracer for CSOA, yielding good correlations ($R^2 > 0.9$) with the time series of CSOA. The correlation of the time series of $C_2H_4N^+$ with CSOA is slightly poorer ($0.83 < R^2 < 0.9$). Since also the difference of the contribution of the ion $C_2H_4N^+$ to the mass spectra of CSOA compared to the mass spectra of other factors is less pronounced than for the ion $C_5H_{10}N^+$ the latter one seems to be more specific for cigarette emissions than $C_2H_4N^+$. This is why the use of $C_5H_{10}N^+$ as a tracer ion for cigarette emissions was favoured over the ion $C_2H_4N^+$.

The CSOA mass spectra from both POPE campaigns show reasonable to very good agreement with CSOA mass spectra reported by Faber et al. (2013) ($0.65<R^2<0.96$). Some differences are observed between the mass spectra obtained for POPE2013 and 2014 in the relative fraction of CO_2^+ and its related ions, which also affects the observed elemental ratios (Fig. 12). This results in a potential error in CSOA concentrations of less than 10 % and probably is due to a PMF artefact and/or insufficient correction for gas phase CO_2. Comparison with mass spectra of cigarette smoke obtained in the laboratory (Faber et al., 2013) show a contribution of CO_2^+ more similar to the POPE2013 measurements. Further laboratory work in order to obtain more robust source spectra is needed to better constrain the expected f44 in CSOA mass spectra. Due to the low intensity of the N-containing ions and a conservative selection of ions which were fitted in the mass spectra reported by Faber et al. (2013), the nicotine fragment ($C_5H_{10}N^+$) was not observed in their measurements of cigarette smoke. However, after re-analysis of the mass spectra with integration of the $C_5H_{10}N^+$ ion in the fitting procedure, a contribution of the nicotine tracer ion is clearly visible (Peter Faber, personal communication).

Also the time series of mass concentration of the CSOA factor clearly support its attribution to cigarette smoking emissions. The diurnal cycle of the CSOA factor strongly correlates with typical working hours at the measurement location and with the diurnal cycle of the marker fragment $C_5H_{10}N^+$ ($R^2 = 0.98$ for POPE2013 and POPE2014), exemplarily shown for POPE2014 in Fig. 13 (top). Averaged CSOA mass concentrations for each day of the week (Fig. 13; bottom) show distinct differences between working days and weekend, when the administration of the hospital where the measurements took place was closed, supporting the attribution of this PMF factor to locally emitted CSOA. Very similar observations were made during the POPE2013 measurements.

Particles from cigarette smoke contributed 9-24 % (0.62-0.76 µg m^{-3}) to the total organic aerosol mass measured at the location in central Rome. No indications for cigarette emissions were found during the DIAPASON measurements. This result shows a potentially strong influence on air quality in the direct environment of smokers, like it was also observed by Faber et al. (2013) and Fröhlich et al. (2015). Since in our measurement, CSOA is mostly produced close to the measurement location (i.e. can be regarded a local contamination), it was not included in the previous analyses of organic aerosol composition (see Sect 4.1).

$C_5H_{10}N^+$ as a potential CSOA marker ion: The ion $C_5H_{10}N^+$ was further investigated in order to assess its applicability and limitations as a tracer for cigarette emissions in AMS data sets. While the nicotine fragment ion $C_5H_{10}N^+$ (*m/z* 84.08) seems to be unique to cigarette emissions, the proximity of the ions $C_5H_8O^+$ (*m/z* 84.06) and $C_6H_{12}^+$ (*m/z* 84.09) in the mass spectra causes interferences, since the mass resolution of the instrument ($R \approx 2000$ in V-mode) is not sufficient for completely separating the individual ion signals. Mass spectra of the primary organic aerosol PMF factors (HOA, COA, BBOA) which are not related to cigarette emissions (“POA_{noCSOA}”) show significant contributions (0.2 to 0.7 % of total organics mass spectral signal) of the ions $C_5H_8O^+$ and $C_6H_{12}^+$, leading to an artificial increase of the nicotine tracer ion signal. In contrast, for OOA mass spectra the contribution of these ions is comparatively low (0.1 % of total organics mass spectral signal), which is why a potential interference of OOA was neglected in the following considerations.

AMS measurements performed during DIAPASON2013 and DIAPASON2014, which are assumed to be not influenced by local cigarette emissions, were used to quantify the concentration-dependent influence of POA_{noCSOA} on the nicotine tracer ion $C_5H_{10}N^+$ due to fitting interferences from the neighbour ions $C_5H_8O^+$ and $C_6H_{12}^+$. A linear relationship between $C_5H_{10}N^+$ and POA_{noCSOA} was observed for a POA_{noCSOA} concentration range of 0-10 µg m^{-3}. Using the mass contribution of $C_5H_{10}N^+$ to the total CSOA mass spectra (Fig. 12; 1.8 % and 1.9 % for POPE2013/2014, respectively), the corresponding ion signals were converted into CSOA detection limits. It was found that, under conservative considerations, CSOA concentrations of at least 10 % of POA_{noCSOA} are needed in order to exceed detection limits. During conditions of negligible POA_{noCSOA} concentrations, a CSOA detection limit of 80 ng m^{-3} was estimated.

The same considerations were performed for the ion $C_2H_4N^+$ in order to evaluate its applicability as a nicotine tracer. In contrast to $C_5H_{10}N^+$, the ion $C_2H_4N^+$ seems to be less affected by its neighbouring ions. Therefore we assume that its occurrence in the mass spectra of different PMF factors does not only result from interferences of signals, but that $C_2H_4N^+$ is also included in non-CSOA factors. However, the contribution of $C_2H_4N^+$ to the CSOA mass spectra is higher than its contribution to those of other aerosol species (1-2 % instead of 0.1-0.6 %). Therefore, we would recommend using $C_2H_4N^+$ as a second nicotine tracer, but only in combination with $C_5H_{10}N^+$.

Based on these estimations, it can be concluded that $C_5H_{10}N^+$ is a suitable nicotine tracer ion for HR-AMS measurements which are influenced by local cigarette emissions (i.e. CSOA larger than 10 % of POA_{noCSOA}) and can be used to estimate CSOA concentrations or to identify a CSOA factor from PMF analysis. Urban background concentrations of cigarette-related particles in the range of 1 % of PM_1, as reported by Rogge et al. (1994), however, are below the estimated CSOA detection limits. This is still true when considering typical contributions of OOA to total organics (~50 %) and of OA to PM_1 (also ~50 %), leading to a CSOA detection limit of around 2.5 % of PM_1. In order to identify cigarette smoke contributions in the order of 1 % of PM_1, higher mass spectral resolution is needed to be able to separate the nicotine tracer ion from its neighbouring ions. Nevertheless, the fitting of $C_5H_{10}N^+$ at *m/z* 84 and also $C_2H_4N^+$ at *m/z* 42 could be worthwhile for HR-AMS data sets which are potentially influenced by cigarette-related particles. In future work, it should be investigated how aging processes affect the appearance of the marker ion $C_5H_{10}N^+$ in the mass spectra.

5 Summary and conclusions

Intensive field campaigns have been performed during Oct/Nov 2013 and May/June 2014, each time consecutively at two locations (suburban, urban) in the area of Rome, enabling the study of seasonal and spatial differences of aerosol and trace gas characteristics.

During both years at the suburban location an impact of advected Saharan dust on PM_{10} levels was detected. With respect to background conditions, increases of PM_{10} by 150 % (2013) and 100 % (2014) were measured, corresponding to average absolute increases of about 12-17 µg m^{-3}. No influence of the dust occurrence on the non-refractory PM_1 chemical composition was found during the advections.

At both locations, during the Oct/Nov measurements air quality was more strongly influenced by primary emissions (e.g. BC, NO_x, PAH) with generally increased particle number concentration (PNC), whereas during May/June the contribution from secondary particles (sulphate, aged OA) and ozone was more important. Also during May/June 2014, new particle formation was frequently detected around midday, while in the colder season no distinct increase of PNC took place outside typical rush hour times. The consequence of higher temperatures plus stronger solar radiation was also visible in the SOA-related fraction of the organic aerosol: During the warmer season two types of OOA (less oxidised, fresher SV-OOA and strongly oxidised, older LV-OOA) were identified, while during the colder season only strongly oxidised OOA was found.

Typical tracers for anthropogenic emissions (CO_2, NO_x, PAH, HOA, COA) were increased at the urban measurement location. However, absolute concentrations of POA were higher at the suburban location, due to a strong contribution from biomass burning OA, which here accounted for 1.28 µg m^{-3} (24 % of total OA) and for 0.82 µg m^{-3} (18 % of total OA) in Oct/Nov 2013 and May/June 2014, respectively. To a large degree this was related to agricultural waste burning in the surrounding areas and during Oct/Nov2013 potentially also to residential wood burning in the urban periphery.

Cooking- and traffic-related aerosol was observed at both locations during both seasons. The diurnal cycles of HOA (traffic-related OA) always peaked during rush hour times. A time shift in the morning rush hour peak between the suburban site and central Rome was observed, likely as a consequence of traffic progressing from the suburbs to the city centre. HOA accounted for 0.59 to 0.93 µg m^{-3} (13 to 29 % of OA) at the different locations and seasons. COA, as an indicator for cooking activities, showed maximum concentrations during lunch and dinner times at central Rome, whereas at the suburban location only during dinner times distinct peaks were observed. Average COA mass concentrations of 0.45 to 0.70 µg m^{-3} (8 to 29 % of OA) were found, with higher concentrations observed at central Rome compared to the suburban location (2013: +50 %; 2014: +23 %), as expected due to the higher density of related sources.

A type of OA related to nearby cigarette emissions (CSOA) was detected at central Rome, and found to strongly correlate with a characteristic nicotine fragment (*N*-methylpyrrolidine, $C_5H_{10}N^+$) at *m/z* 84 in the mass spectra. This ion could serve as a suitable tracer for locally emitted cigarette smoke also for other datasets. However, in order to identify CSOA based solely on this tracer ion, CSOA must account for at least 10 % of the sum of COA, HOA and BBOA, due to interferences of neighbouring ion signals from these POA types. In the absence of those, a detection limit of 80 ng m^{-3} was found for CSOA. These findings imply that the resolution of the HR-AMS is not sufficient to identify urban background contributions of cigarette emissions (~1 % of PM_1, Rogge et al., 1994) based solely on $C_5H_{10}N^+$, while fitting of this ion could be worthwhile for HR-AMS datasets which are potentially influenced by nearby cigarette emissions.

During our measurements sub-micron aerosol originating from sources in the metropolitan area of Rome and particles being advected from outside (dust periods were excluded) contributed 42-70 % and 30-58 % to total measured PM_1, respectively. Thus, during our measurements approximately half of the locally measured PM_1 was “home-made”.

While for individual aerosol types clear spatial and temporal characteristics were observed and can be understood, no general conclusion can be drawn whether total aerosol mass concentrations are generally higher at the suburb or the city centre. Instead, consistent with observations made in the area of Paris (Freutel et al., 2013), it was found that aerosol levels strongly depend on the combination of meteorological conditions (e.g. origin of air masses, dilution capacity within the boundary layer) and contributions of secondary aerosols and local emissions.

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Table 1. Summary of measurement campaigns and measurement periods.

Campaign name	Measurement location	Classification	Measurement period
DIAPASON2013	Tor Vergata	Urban background	23.10. – 07.11.2013
POPE2013	Central Rome	Urban	07.11. – 14.11.2013
DIAPASON2014	Tor Vergata	Urban background	20.05. – 04.06.2014
POPE2014	Central Rome	Urban	04.06. – 17.06.2014

Table 2. Summary of the instruments deployed in MoLa during all measurement periods, together with measured variables. D_p: particle diameter (defined according to individual instrumental measurement method: optical, aerodynamic, or mobility diameter).

Instrument	Measured variable
Chemical composition of PM_1	
Aerosol mass spectrometer (HR-ToF-AMS)	Organics, sulphate, nitrate, ammonium, chloride (non-refractory) mass concentrations (D_p = ~ 70 – 800 nm)
Multi Angle Absorption Photometer (MAAP)	Black carbon (BC) mass concentration
PAS2000	Particulate PAH[a] mass concentration
Physical aerosol properties	
Condensation particle counter (CPC)	Total number concentration (D_p > 2.5 nm)
Fast mobility particle sizer (FMPS)	Size distribution (D_p = 5.6 – 560 nm)
Optical particle counter (OPC)[b]	Size distribution (D_p = 0.25 – 32 µm)
Aerodynamic particle sizer (APS)	Size distribution (D_p = 0.5 – 20 µm)
Environmental Dust Monitor (EDM)	Mass concentration of PM_1, $PM_{2.5}$, PM_{10}
Trace gas mixing ratios	
Airpointer	NO_2, NO_x, NO, SO_2, CO, O_3
LI-840	H_2O, CO_2
Meteorology	
Meteorological station[c]	Wind direction, wind speed, temperature, pressure, solar radiation, precipitation, relative humidity

[a] Polycyclic Aromatic Hydrocarbons

[b] No data collected during DIAPASON2014.

[c] Wind direction, wind speed and solar radiation data not useable during POPE measurements.

Table 3. Summary of selected variables measured during DIAPASON (Tor Vergata) and POPE (central Rome) in 2013 and 2014. Values represent total campaign averages, calculated from 5 min averages, and their standard deviations. N/A: not available.

	DIAPASON2013 (Oct/Nov 2013)	POPE2013 (Nov 2013)	DIAPASON2014 (May/June 2014)	POPE2014 (June 2014)
Temperature / °C	17.6 ± 2.9	16.1 ± 2.8	19.0 ± 3.5	24.9 ± 4.5
Rain [a] / mm	46.3 (5)	25.9 (6)	6.3 (5)	69.9 (4)
Pressure / hPa	1004 ± 7	1009 ± 6	1002 ± 2	1012 ± 4
Solar radiation [b] / W m^{-2}	103 ± 26	N/A	282 ± 55	N/A
Rel. humidity / %	78 ± 11	73 ± 6	61 ± 15	52 ± 16
Wind speed / m s^{-1}	1.9 ± 1.3	N/A	2.5 ± 1.6	N/A
Daily BLH [c] max. / m	850 ± 220	900 ± 150	1500 ± 450	1560 ± 250
TKE [d] / J kg^{-1}	0.48 ± 0.45	1.8 ± 2.3	0.84 ± 0.79	0.68 ± 0.52
PM_{10} [e] / µg m^{-3}	22 ± 12	13 ± 7	15 ± 8	17 ± 8
$PM_{10-2.5}$ [e,f] / µg m^{-3}	4.5 ± 3.1	4.7 ± 3.2	9.3 ± 6.5	7.1 ± 5.7
PM_1 [e] / µg m^{-3}	15 ± 10	6.0 ± 3.6	5.8 ± 3.5	7.5 ± 3.8
PM_1 (AMS+BC) [g,i] / µg m^{-3}	12 ± 8	5.6 ± 4	8.7 ± 5	12.5 ± 6
PNC [h] / 10^3 cm^{-3}	23 ± 15	27 ± 11	18 ± 10	13 ± 5
NO_x / ppb	29 ± 27	36 ± 27	9 ± 9	13 ± 8
CO_2 / ppm	410 ± 20	420 ± 20	410 ± 20	420 ± 20
O_3 / ppb	14 ± 14	8.9 ± 10	34 ± 19	35 ± 22
PAH / ng m^{-3}	45 ± 54	45 ± 39	10 ± 14	12 ± 11
Org [i] / µg m^{-3}	5.3 ± 4.5	2.5 ± 1.8	4.5 ± 3.2	6.6 ± 3.3

SO_4 / µg m^{-3}	2.0 ± 1.1	0.48 ± 0.44	1.6 ± 0.44	2.6 ± 1.3
NO_3 / µg m^{-3}	0.86 ± 0.80	0.23 ± 0.17	0.61 ± 0.72	0.49 ± 0.40
NH_4 / µg m^{-3}	0.88 ± 0.48	0.20 ± 0.18	0.66 ± 0.27	0.97 ± 0.46
Chl / µg m^{-3}	0.09 ± 0.16	0.04 ± 0.06	0.06 ± 0.13	0.03 ± 0.08
BC / µg m^{-3}	2.9 ± 2.5	2.2 ± 1.7	1.3 ± 1.0	1.8 ± 1.0

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[a] Total accumulated amount of rain during measurements. Numbers of days with rain are given in parentheses.

[b] Average and standard deviation of daily means. Only days with 24 h measurements were used (includes 85 % of data).

[c] Average and standard deviation of daily boundary layer height (BLH) maxima from polarization LIDAR-ceilometer. measurements at the Tor Vergata site provided by the DIAPASON project (DIAPASON, 2016).

[d] Turbulent kinetic energy (TKE) calculated from 2-hour averages provided by Arpa Lazio Environmental Agency from measurements at 3 sites around Rome (Tor Vergata, Castel di Guido, Boncompagni).

[e] PM concentrations from EDM measurements.

[f] Difference between PM_{10} and $PM_{2.5}$ (coarse particles).

[g] PM_1 concentration based on the sum of AMS species and BC.

[h] PNC from CPC measurements.

[i] For POPE2013 and POPE2014 corrected for contribution from local cigarette smoke emissions, compare Sect. 4.2.5.

Table 4: Estimated contribution of "home-made" and “advected” species to total PM_1 for all measurement periods. Dust advection periods and emissions from cigarettes are excluded.

	Home-made PM_1 / µg m^{-3} (contribution to PM_1 / %)	Advected PM_1 / µg m^{-3} (contribution to PM_1 / %)
DIAPASON2013	6.5 (47)	7.3 (53)
POPE2013	3.9 (70)	1.7 (30)
DIAPASON2014	4.7 (59)	3.3 (41)
POPE2014	5.2 (42)	7.3 (58)

Table 5. Summary of “dust” and “no dust” periods identified during DIAPASON2013 and DIAPASON2014, including mean values and standard deviation of $PM_{10-2.5}$, PM_{10} and PM_1 (all from EDM measurements).

	DIAPASON2013		**DIAPASON2014**	
	Dust	No dust	Dust	No dust
Period	23.10. - 01.11.13	02.11. - 07.11.13	20.05. - 26.05.14	27.05. - 04.06.14
Mean $PM_{10-2.5}$ / µg m^{-3}	5.4 ± 3.1	3.0 ± 2.5	15.4 ± 5.9	5.5 ± 2.8
Mean PM_{10} / µg m^{-3}	28 ± 10	11 ± 6	22 ± 7	11 ± 5
Mean PM_1 / µg m^{-3}	20 ± 9 (71% of PM_{10})	6.3 ± 4.1 (57% of PM_{10})	7.0 ± 3.3 (32% of PM_{10})	5.0 ± 3.5 (45% of PM_{10})

Table 6. Variables measured during DIAPASON2014 indicating different ambient conditions on days classified as "nucleation", "non-defined" and "non-nucleation" days. Total particle number concentrations (PNC) and the classification criteria and number of respectively classified days are also listed. For each measurement day and presented variable daily averages for the period 10 am – 4 pm were calculated. Here, mean and standard deviation of these averages are shown with maximum values coloured in red, minimum values coloured in green.

	Nucleation days	Non-defined days	Non-nucleation days
Classification	$PNC_{nuc}/PNC_{bg} \geq 1.5$	$1 < PNC_{nuc}/PNC_{bg} < 1.5$	$PNC_{nuc}/PNC_{bg} \leq 1$
Number of days	6	6	2
Rain [a] / mm	0 (0)	0.7 (1)	2.3 (1)
Total PNC / 10^3 cm^{-3}	25 ± 9	13 ± 3	12 ± 1
Temperature / ° C	23 ± 2	22 ± 2	20 ± 0
Rel. humidity / %	43 ± 6	46 ± 9	56 ± 4
Solar radiation / W m^{-2}	780 ± 80	700 ± 190	670 ± 200
O_3 / ppb	54 ± 4	47 ± 6	48 ± 1
Total PM_1 organics / µg m^{-3}	3.9 ± 2.4	3.6 ± 1.0	2.9 ± 2.3

[a] Total accumulated amount of rain (10 am – 4 pm). Numbers of days with rain are given in parentheses.

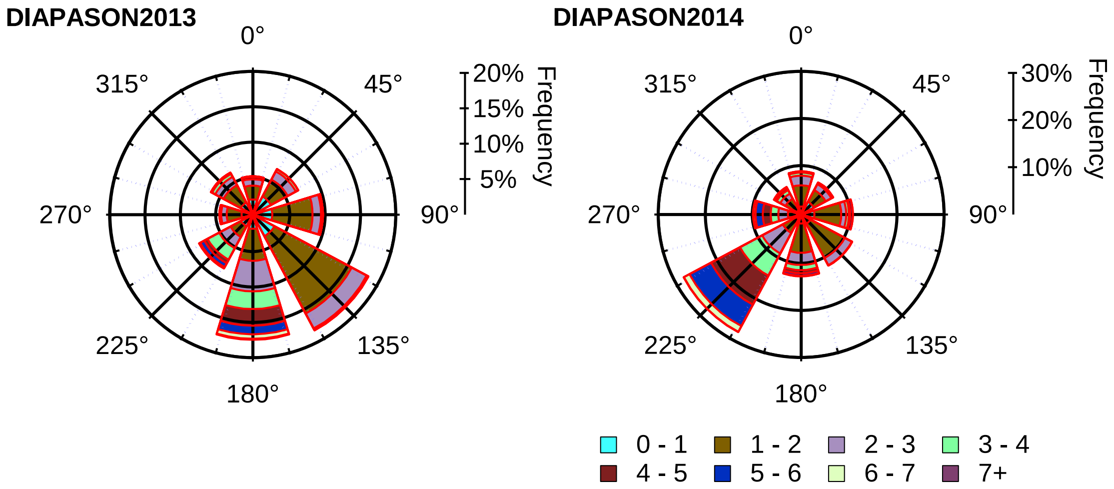

Figure 1. Relative frequency of local wind directions (in °) colour coded with wind speed measured during DIAPASON2013 (left) and DIAPASON2014 (right).

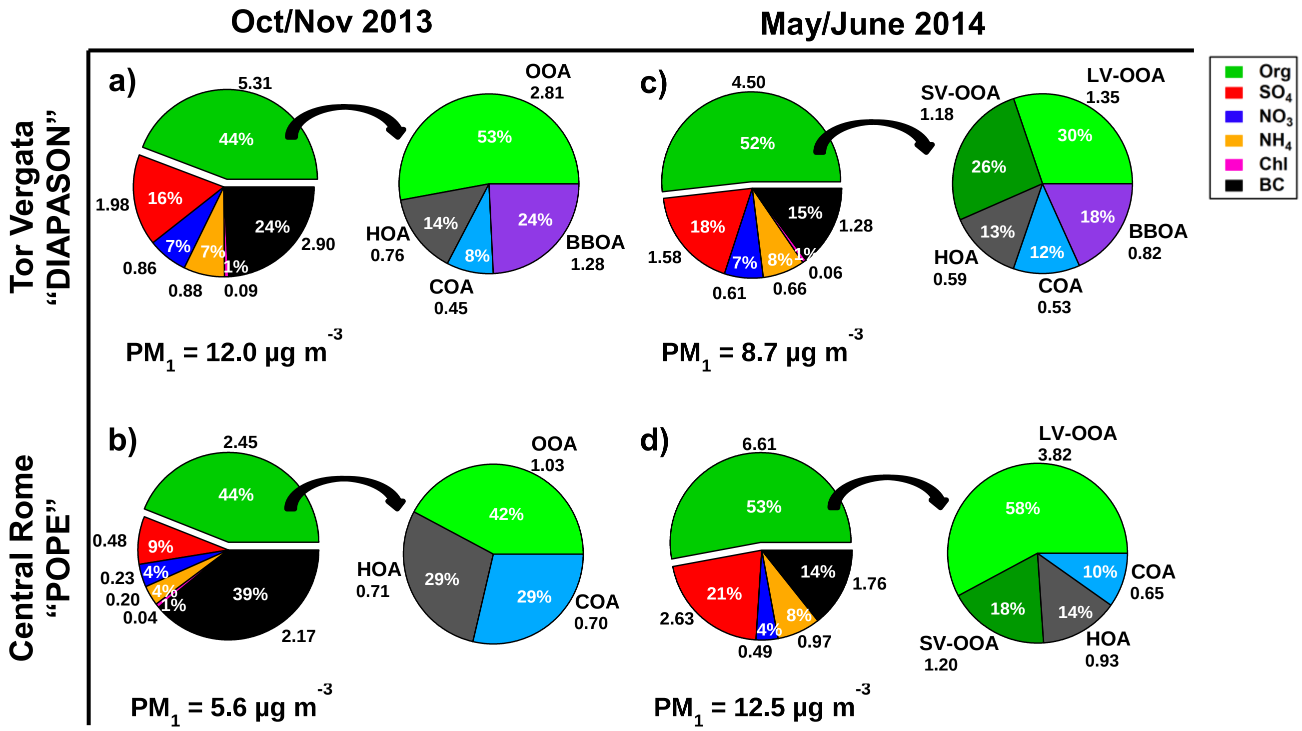

Figure 2. Mean chemical composition (μg m^{-3}) of non-refractory PM$_1$ together with BC (left chart in panels a-d) and PMF-separated organic fraction (right chart in panels a-d) for each measurement period. PM$_1$ values below the pie charts represent total mass concentration of AMS-measured species plus BC. The organic fraction measured at central Rome was corrected for contributions from cigarette smoke in the local environment (Sect. 4.2.5, also omitted from the pie charts depicting the PMF-separated organic fraction).

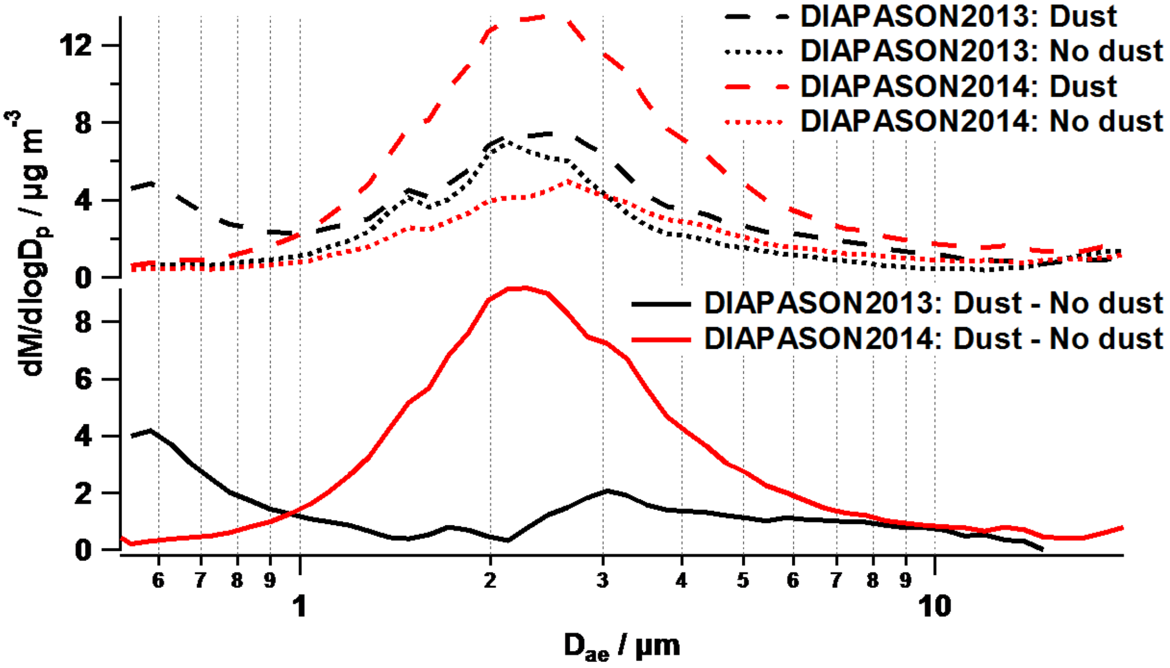

Figure 3. Size-resolved mass distribution (dM/dlogD$_p$) during “dust” (dashed traces) and “no dust” (dotted traces) periods (top panel) measured with the APS during DIAPASON2013 (black traces) and DIAPASON2014 (red traces). The difference of the size-resolved mass distributions measured during “dust” and “no dust” periods indicate the size distributions of the dust particles measured during both years (bottom panel). D$_{ae}$ is the aerodynamic particle diameter.

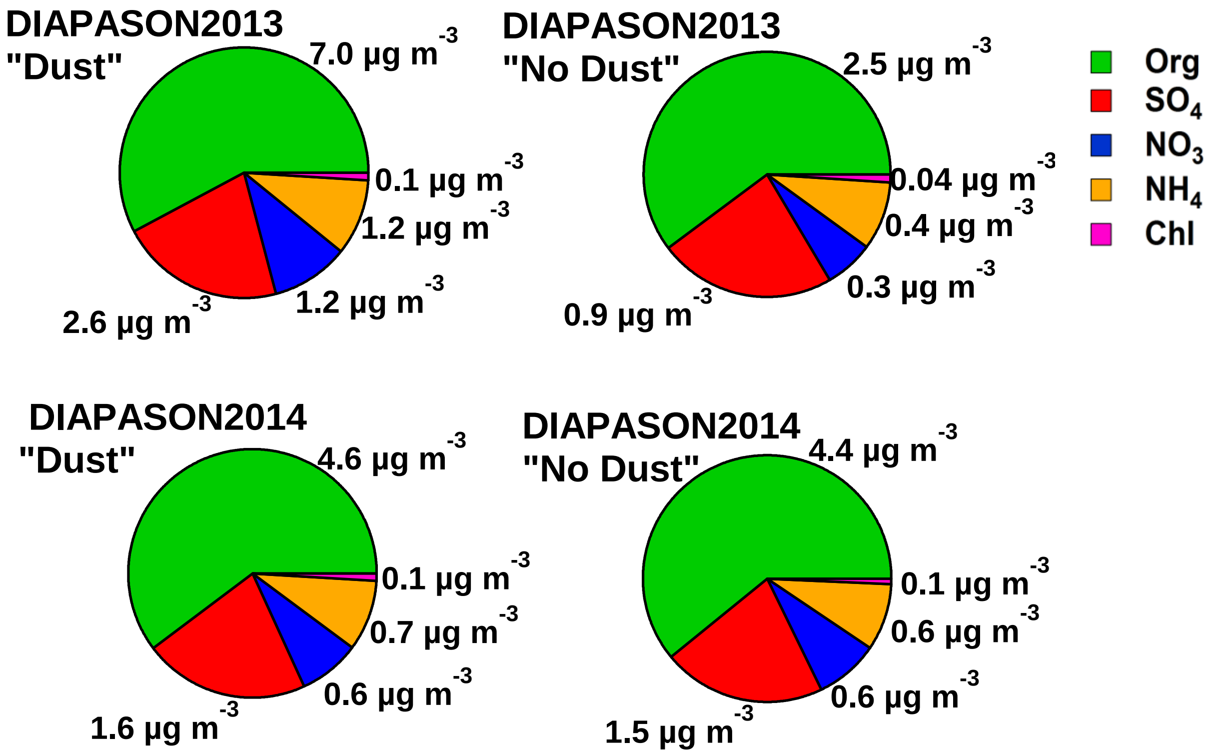

Figure 4. Comparison of mean chemical composition of non-refractory PM_1 obtained from AMS measurements during “dust” (left) and “no dust” (right) periods during DIAPASON2013 (top) and DIAPASON2014 (bottom).

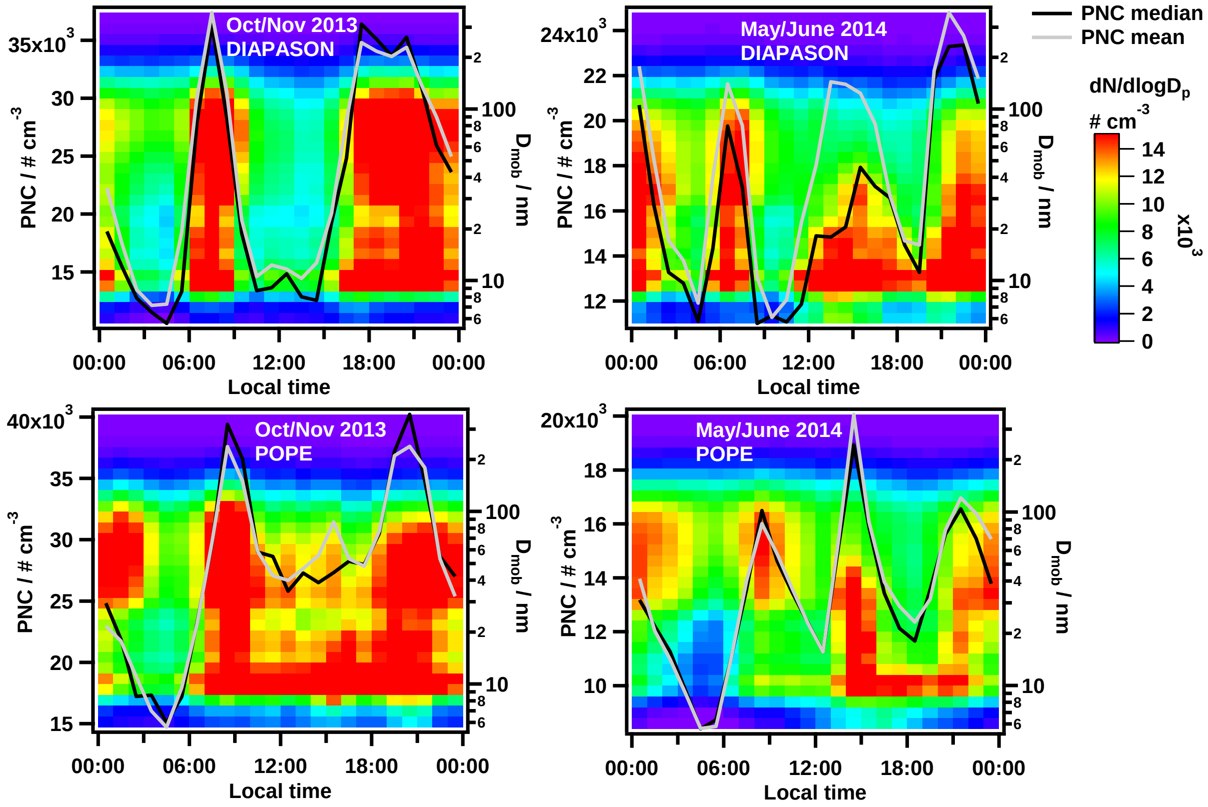

Figure 5. Average diurnal cycles of particle number concentrations and size distributions for DIAPASON (top) and POPE (bottom) for each year (left: 2013, right: 2014). Image plots of diurnal cycles of the particle number size distributions (colour coded for dN/dlogD$_p$) are shown with the particle diameters on the right axes (mobility particle diameter D$_{mob}$). Mean (grey) and median (black) diurnal cycles of the total particle number concentrations are shown on the left axes. New particle formation at midday was only observed in the May/June 2014 campaigns.

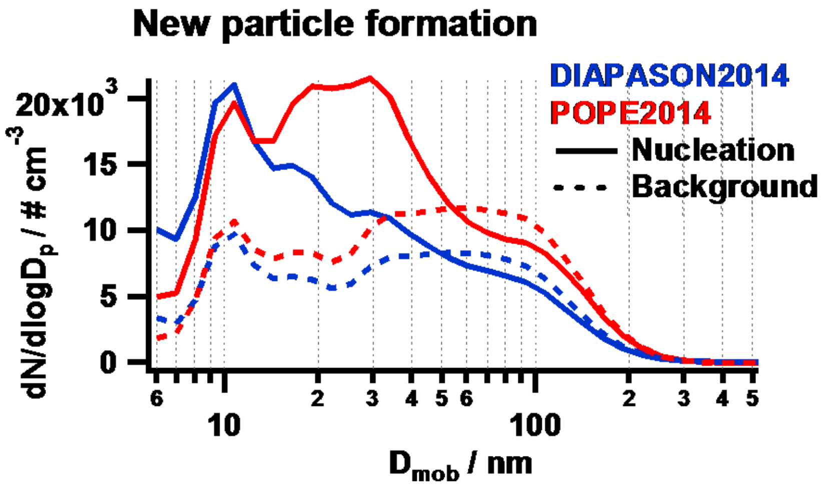

Figure 6. Average size distributions of particle number concentrations (dN/dlogD$_p$) for maximum PNC at midday (DIAPASON2014: 1-3 pm; POPE2014: 2-3 pm; solid traces) and during background conditions (10-11 am; dashed traces) from FMPS measurements for DIAPASON2014 (blue) and POPE2014 (red). D$_{mob}$ is the mobility diameter.

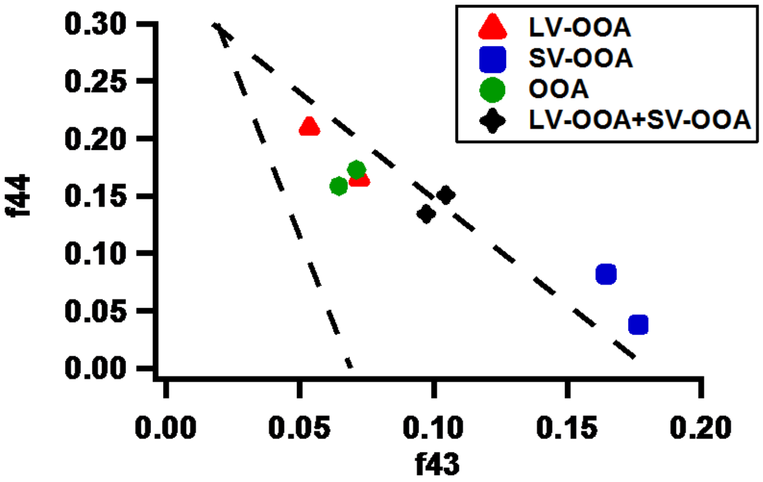

Figure 7. f44 vs. f43 plotted for each OOA factor obtained from PMF analysis of the organic fraction of HR-AMS data. OOA factors (green markers) resulted from DIAPASON2013 and POPE2013 measurements; SV-OOA (blue markers) and LV-OOA factors (red markers) were found during DIAPASON2014 and POPE2014 measurements. The recombination of the factors LV-OOA and SV-OOA for both DIAPASON2014 and POPE2014 is also shown (black markers). The dashed lines represent the triangular space in which measured ambient OOA components typically cluster according to Ng et al. (2010).

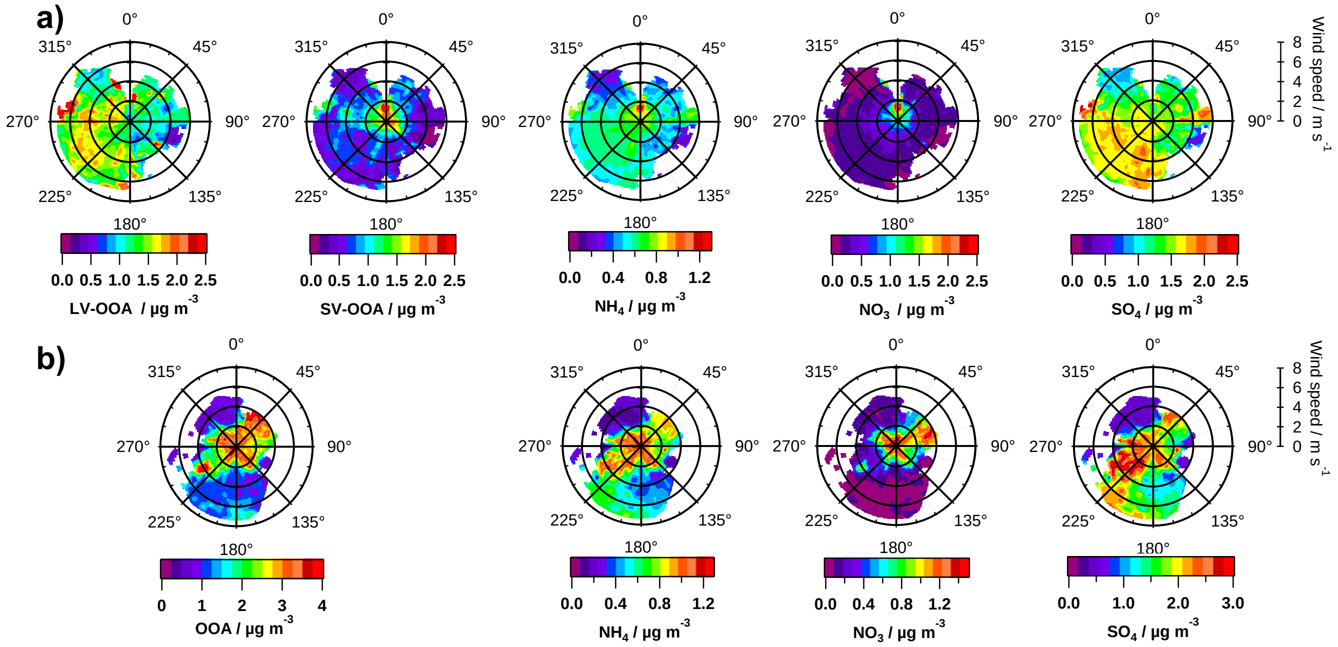

Figure 8. Polar plots of (a) submicron LV-OOA, SV-OOA, NH_4, NO_3 and SO_4 concentration (colour coded) obtained from DIAPASON2014 measurements and (b) submicron OOA, NH_4, NO_3 and SO_4 concentration obtained from DIAPASON2013 measurements as a function of local wind direction (°) and speed (m s^{-1}).

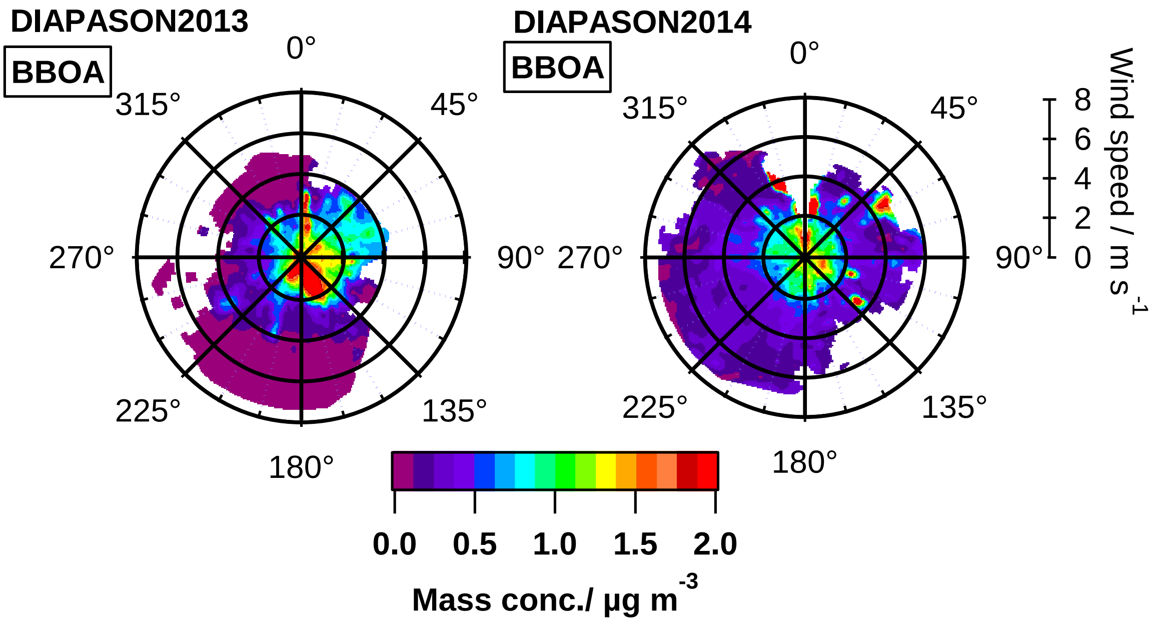

Figure 9. Submicron BBOA mass concentrations (colour coded) as a function of local wind direction (°) and speed (m s^{-1}) for DIAPASON2013 (left) and DIAPASON2014 (right).

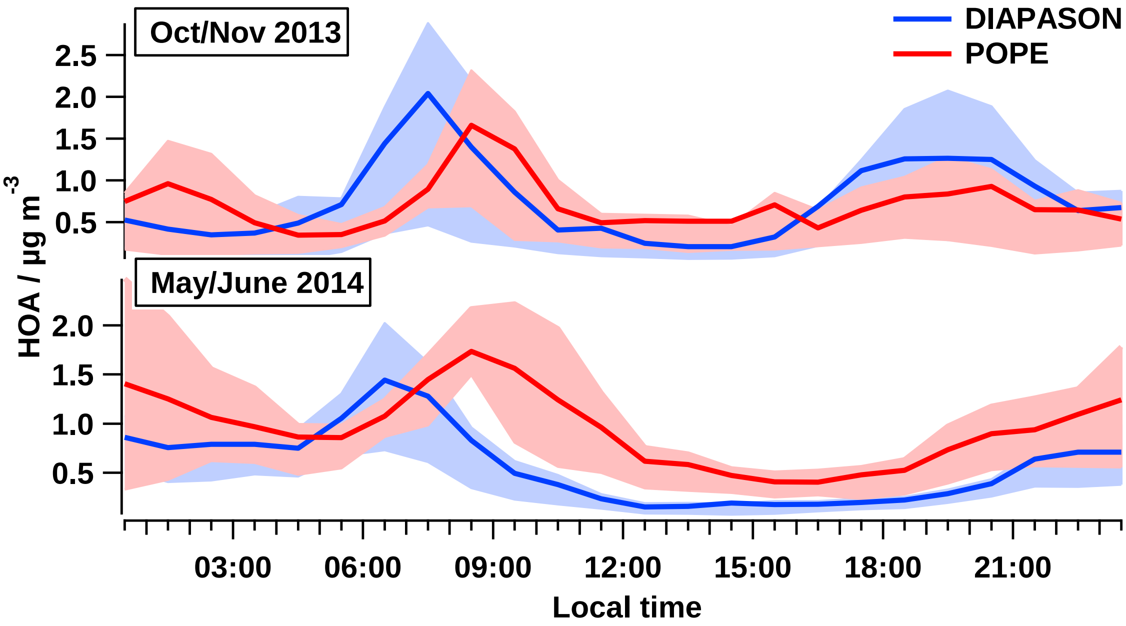

Figure 10. Diurnal cycles of HOA mass concentrations obtained from measurements at Tor Vergata (blue; DIAPASON) and central Rome (red; POPE) during both seasons. Shown are mean concentrations (traces) and the corresponding 25th and 75th percentiles (shaded areas).

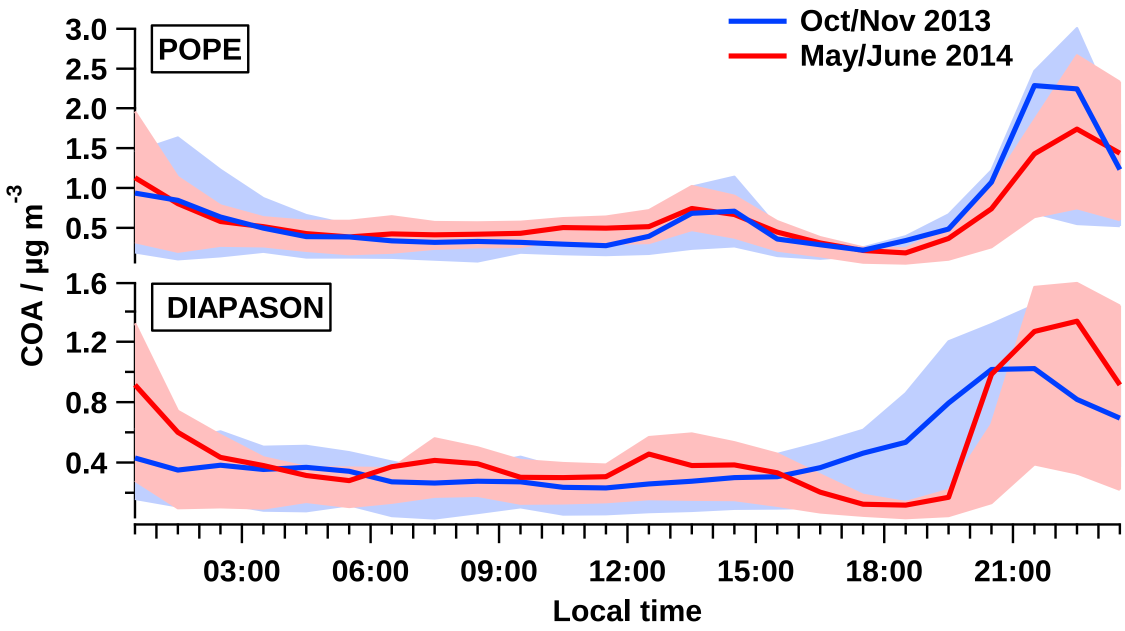

Figure 11. Diurnal cycles of COA mass concentrations observed at central Rome (top; POPE) and the suburban site (bottom; DIAPASON) during both seasons. Shown are mean concentrations (traces) and the corresponding 25th and 75th percentiles (shaded areas).

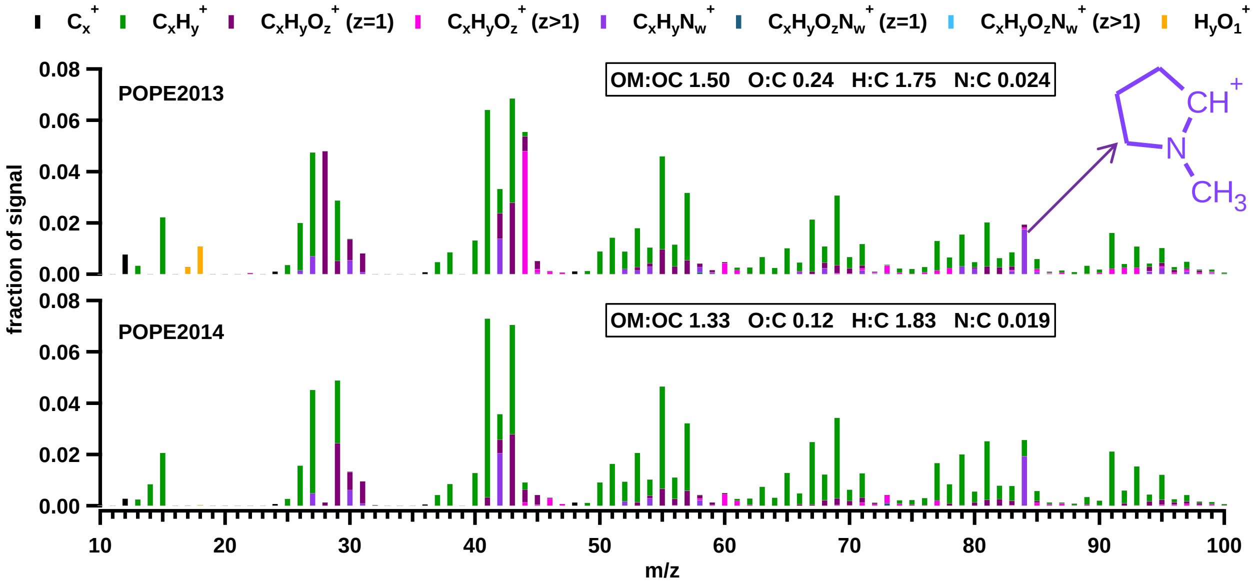

Figure 12. Unit mass resolution spectra of CSOA obtained for the two POPE campaigns, calculated from organic high resolution mass spectra and colour coded for the different groups of ion fragments. The elemental ratios are shown in boxes. The chemical structure of the suggested ion fragment at *m/z* 84 ($C_5H_{10}N^+$) is also illustrated.

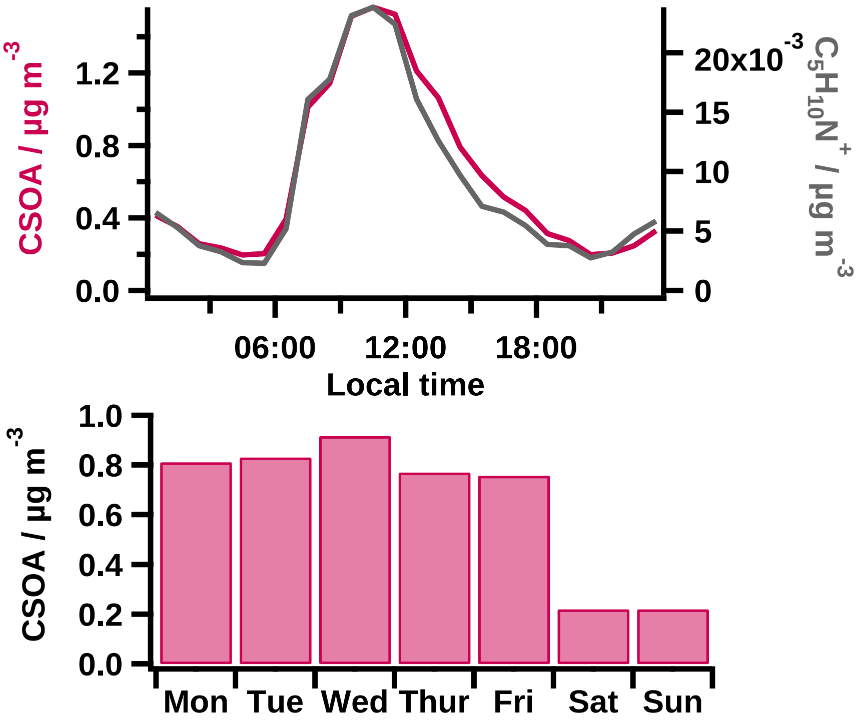

Figure 13. Diurnal cycle of CSOA and the marker fragment $C_5H_{10}N^+$ (top) and the weekly cycle (bottom) of CSOA mass concentrations obtained from HR-AMS data of POPE2014.