



1 Molecular distributions and stable carbon isotope compositions
2 of oxalic acid and related SOA in Beijing before, during and
3 after the 2014 APEC
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43 **Abstract:** To ensure the good air quality for the 2014 APEC, stringent emission controls were
44 implemented in Beijing and its surrounding regions, leading to a significant reduction in $PM_{2.5}$
45 loadings. To investigate the impacts of the emission controls on aerosol composition and
46 formation, high-volume $PM_{2.5}$ samples were collected in Beijing from 08/10/2014 to 24/11/2014
47 and determined for secondary inorganic ions (SIA, i.e., SO_4^{2-} , NO_3^- and NH_4^+), dicarboxylic
48 acids, keto-carboxylic acid and α -dicarbonyls, as well as stable carbon isotope composition of
49 oxalic acid (C_2). Our results showed that SIA in $PM_{2.5}$ are 52 ± 47 , 18 ± 13 and $33 \pm 29 \mu g m^{-3}$
50 before-, during- and after-APEC, accounting for 29%, 18% and 20% of $PM_{2.5}$, respectively. As
51 the leading dicarboxylic acid, C_2 in $PM_{2.5}$ during the three phases are 502 ± 564 , 101 ± 69 and
52 $166 \pm 157 ng m^{-3}$, accounting for 46%, 31% and 34% of total detected organic compounds
53 (TDOC, i.e., the sum of dicarboxylic acids, keto-carboxylic acids and α -dicarbonyls). The higher
54 values of concentrations and relative abundances of SIA and C_2 before-APEC suggest that $PM_{2.5}$
55 aerosols during this period are more enriched with secondary products, mainly due to an
56 enhanced photochemical oxidation under the higher temperature and more humid conditions.
57 SIA, C_2 and related SOA in $PM_{2.5}$ during-APEC were 2–4 times lower than those before-APEC.
58 C_2 in the regional air masses, which mostly occurred before-APEC, are abundant and
59 enriched in ^{13}C . On the contrary, C_2 in the long-range transport air masses, which mostly
60 occurred during-APEC, is much less abundant but still enriched in ^{13}C . In the local air masses,
61 which mostly occurred after-APEC, C_2 concentration is lower than that before-APEC but higher
62 than that during-APEC and enriched in lighter ^{12}C . A comparison on chemical composition of
63 $PM_{2.5}$ and $\delta^{13}C$ values of C_2 in two events that are characterized by the highest $PM_{2.5}$ levels
64 before- and after-APEC, respectively, further showed that after-APEC SIA and TDOC are much



65 less abundant and fine aerosols are enriched with primary organics and relatively fresh,
66 compared with those before-APEC. Such reduction in secondary aerosols after-APEC, along
67 with a similar reduction during-APEC, is largely due to the decreasing temperatures. Our results
68 indicate that the significant reduction in $PM_{2.5}$ during-APEC is mainly due to the efficient
69 emission controls, but the effect of the decreasing temperatures, which suppressed secondary
70 aerosol production, may also take an important role.

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72 **Key words:** Secondary organic aerosols; Emission controls; Sources and formation mechanisms;
73 Aqueous-phase oxidation; Aerosol acidity and water content.

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75 1. Introduction

76 Atmospheric aerosols profoundly impact the global climate directly by scattering and
77 absorbing solar radiation and indirectly by affecting cloud formation and distribution via acting
78 as cloud condensation nuclei (CCN) and ice nuclei (IN). Moreover, atmospheric aerosols exert
79 negative effects on human health because of their toxicity. Due to fast urbanization and
80 industrialization, high level of atmospheric fine particle ($PM_{2.5}$) pollution has been a persistent
81 problem in many cities of China since the nineties of last century (van Donkelaar et al., 2010).
82 As the capital of China and one of the largest megacities in the world, Beijing has suffered from
83 frequent severe haze pollution especially in winter, affecting more than 21 million people by the
84 end of 2014 (Beijing Municipal Bureau of Statistics, 2015) and causing billions of economic
85 loses (Mu and Zhang, 2013). To improve the air quality Beijing government has put many efforts
86 to reduce the pollutant emissions (i.e., SO_2 , NO_x , dust, and volatile organic compounds (VOCs))
87 from a variety of sources.

88 The 2014 Asia-Pacific Economic Cooperation (APEC) summit was hosted in Beijing from
89 the 5th to 11th November. To ensure good air quality for the summit, a joint strict emission control
90 program was conducted from 3rd November 2014 in Beijing and its neighboring provinces
91 including Inner Mongolia, Shanxi, Hebei and Shandong provinces. During this period thousands
92 of factories and power plants with high emissions were shut down and/or halted, all the
93 construction activities were stopped and the numbers of on-road vehicles were reduced. These
94 strict emission controls resulted in the air quality of Beijing during the APEC period being
95 significantly improved, leading to a decrease in $PM_{2.5}$ concentration by 59.2% and an increase in
96 visibility by 70.2% in Beijing during the summit compared with those before the APEC (Tang et



97 al., 2015) and a term of “APEC-Blue” being created to refer to the good air quality. Such strong
98 artificial intervening not only reduced PM_{2.5} and its precursors’ loadings in Beijing and its
99 surrounding areas but also affected the composition and formation mechanisms of the fine
100 particles (Sun et al., 2016).

101 A number of field measurements have showed that particle compositions in Beijing during
102 wintertime haze periods are dominated by secondary aerosols (Guo et a., 2014; Huang et al.,
103 2014; Xu et al., 2015). Rapid accumulation of particle mass in Beijing during haze formation
104 process is often accompanied by continuous particle size growth (Guo et al, 2014; Zhang et al.,
105 2015), which is in part due to the coating of secondary organic aerosols (SOA) on pre-existing
106 particles (Li et al., 2010). Several studies have found that SOA production during the 2014
107 Beijing APEC periods significantly reduced and ascribed this reduction to the efficient regional
108 emission control (Sun et al., 2016; Xu et al, 2015). However, up to now information of the SOA
109 decrease on a molecular level has not been reported. Dicarboxylic acids are the major class of
110 SOA species in the atmosphere and ubiquitously found from the ground surface to the free
111 troposphere (Fu et al., 2008; Myriokefalitakis et al., 2011; Sorooshian et al., 2007; Sullivan et al.,
112 2007). In the current work we measured molecular distributions of dicarboxylic acids,
113 keto-carboxylic acids and α -dicarbonyls and stable carbon isotope composition of oxalic acid in
114 PM_{2.5} aerosols collected in Beijing before, during and after the APEC event in order to explore
115 the impact of the APEC emission control on SOA in Beijing. We first investigated the changes in
116 concentration and composition of dicarboxylic acids and related compounds during the three
117 periods, then we recognized the difference in stable carbon isotope composition of oxalic acid in
118 different air masses in Beijing during the APEC campaign. Finally we compared the different



119 chemical compositions of PM_{2.5} during two heaviest pollution episodes. Our results indicated
120 that in comparison with those before the event, the decrease in PM_{2.5} levels during APEC 2014
121 was caused not only by the emission controls but also by a decreased SOA production due to the
122 decreasing temperatures.

123 **2. Experimental section**

124 **2.1 Sample collection**

125 PM_{2.5} samples were collected by using a high-volume sampler (Brand, USA) from 8th
126 October to 24th November 2014 on the rooftop of a three-storey building located on the campus
127 of China Research Academy of Environmental Sciences, which is situated in the north part of
128 Beijing and close to the 5th-ring road. All the PM_{2.5} samples were collected onto pre-baked (450
129 °C for 8 h) quartz fiber filters (Whatman 41, USA). The duration of each sample collection is 23
130 hr from 08:00 am of the previous day to 07:00 am of the next day. Field blanks were also
131 collected before and after the campaign by mounting a pre-baked filter onto the sampler for 15
132 min without pumping air. After collection, all the filter samplers were individually sealed in
133 aluminum foil bags and stored in a freezer (-18 °C) prior to analysis. Daily values of SO₂, NO_x
134 and meteorological parameters were cited from the website of Beijing Environmental Protection
135 Agency.

136 **2.2 Sample analysis**

137 **2.2.1 Elemental carbon (EC), organic carbon (OC), water-soluble organic (WSOC) and** 138 **inorganic ions**

139 Detailed methods for the analysis of EC, OC, WSOC and inorganic ions in aerosols were
140 reported elsewhere (Li et al., 2011; Wang et al 2010). Briefly, EC and OC in the PM_{2.5} samples



141 were determined by using DRI Model 2001 Carbon analyzer following the Interagency
142 Monitoring of Protected Visual Environments (IMPROVE) thermal/optical reflectance (TOR)
143 protocol (Chow et al., 2007). WSOC and inorganic ions in the samples were extracted with
144 Milli-Q pure water and measured by using Shimadzu TOC-L CPH analyzer and Dionex-600 ion
145 chromatography, respectively (Li et al. 2011; Wang et al 2010).

146 **2.2.2 Dicarboxylic acids, keto-carboxylic acids and α -dicarbonyls**

147 The method of analyzing PM_{2.5} samples for dicarboxylic acids, ketocarboxylic acids and
148 α -dicarbonyl has been reported elsewhere (Wang et al., 2012; Meng et al., 2014; Cheng et al.,
149 2015). Briefly, one eighth of the filter was extracted with Milli-Q water, concentrated to near
150 dryness, and reacted with 14% BF₃/butanol at 100 °C for 1 h to convert aldehyde group into
151 dibutoxy acetal and carboxyl group into butyl ester. Target compounds in the derivatized samples
152 were identified by GC/MS and quantified by GC-FID (Agilent GC7890A).

153 **2.3. Stable carbon isotope composition of oxalic acid (C₂)**

154 Stable carbon isotope composition ($\delta^{13}\text{C}$) of C₂ was measured using the method developed
155 by Kawamura and Watanabe (2004). Briefly, $\delta^{13}\text{C}$ values of the derivatized samples above were
156 determined by gas chromatography-isotope ratio-mass spectrometry (GC-ir-MS) (Thermo Fisher,
157 Delta V Advantage). The $\delta^{13}\text{C}$ value of C₂ was then calculated from an isotopic mass balance
158 equation based on the measured $\delta^{13}\text{C}$ of the derivatizations and the derivatizing reagent
159 (1-butanol) (Kawamura and Watanabe, 2004). Each sample was measured for three times to
160 ensure the difference of the $\delta^{13}\text{C}$ values less than 0.2‰, and the isotope data reported here is the
161 averaged value of the triplicate measurements.

162 **3. Results and discussion**



163 **3.1 Variations in meteorological conditions, gaseous pollutants and major components of**
164 **PM_{2.5} during the Beijing 2014 APEC campaign**

165 Based on the emission control implementation for the APEC, we divided the whole study
166 period into three phases: before-APEC (08/10 to 02/11), during-APEC (03/11 to 12/11) and
167 after-APEC (13/11 to 24/11). Temporal variations in meteorological parameters and
168 concentrations of gaseous pollutants and major components of PM_{2.5} during the three phases are
169 shown in Fig. 1 and summarized in Table 1.

170 Temperature during the sampling campaign showed a continuous decreasing trend with
171 averages of 13 ± 2.6 °C, 7.0 ± 1.7 °C and 4.3 ± 1.3 °C before-, during- and after-APEC periods,
172 respectively, while relative humidity (RH) did not show a clear trend with mean values of $62 \pm$
173 19% , $47 \pm 14\%$ and $51 \pm 16\%$ during the three periods (Fig. 1a and Table 1). SO₂ showed a
174 similar level before- and during-APEC periods (8.8 ± 4.6 μg m⁻³ versus 7.6 ± 3.9 μg m⁻³) (Table
175 1), but increased dramatically to 23 ± 8.8 μg m⁻³ after-APEC due to domestic coal burning for
176 house heating (Fig. 1b). NO₂ concentration (45 ± 18 μg m⁻³) during the APEC reduced by about
177 30% compared to that in the before- and after-APEC phases (71 ± 27 μg m⁻³ versus 78 ± 29 μg
178 m⁻³) (Table 1), mainly because of the reduction of the on-road vehicle numbers, as well as the
179 reduced productivities of power plant and industry. O₃ displayed a decreasing trend similar to
180 that of temperature (Fig. 1c). PM_{2.5} pollution episodes in Beijing showed a periodic cycle of 4–5
181 days, which is caused by the local weather cycles. Secondary inorganic ions (SIA, i.e., SO₄²⁻,
182 NO₃⁻ and NH₄⁺) are major components of PM_{2.5} and present a temporal variation pattern similar
183 to that of the fine particles (Fig. 1d). In the current work mass ratio of NO₃⁻/SO₄²⁻ in PM_{2.5}
184 during the whole study time is 1.8 ± 1.9 (Table 1), which is in agreement with the ratio (1.6–2.4)



185 for PM_1 observed during the same time by using aerosol mass spectrometry (AMS) (Sun et al.,
186 2016). OC and EC of $PM_{2.5}$ linearly correlated each other ($R^2=0.91$) and varied periodically in a
187 cycle similar to SIA (Fig. 1e). OC/EC ratio during the whole sampling period is 3.3 ± 0.6 (range:
188 2.2–4.7) with no significant differences among the three APEC phases (Table 1).

189 Figure 2 shows the differences in chemical composition of $PM_{2.5}$ before-, during- and
190 after-APEC periods. $PM_{2.5}$ is $98 \pm 46 \mu\text{g m}^{-3}$ during-APEC, about 50% lower than that before-
191 and after-APEC ($178 \pm 122 \mu\text{g m}^{-3}$ versus $161 \pm 100 \mu\text{g m}^{-3}$), respectively. Organic matter (OM)
192 is the most abundant component of the fine particles. Relative abundance of OM (OM, 1.6 times
193 of OC) (Xing et al., 2013) to $PM_{2.5}$ increase from 24% before-APEC to 30% and 39% during-
194 and after-APEC, although the mass concentration ($19 \pm 7.6 \mu\text{g m}^{-3}$) of OC during-APEC is the
195 lowest compared to those before- and after-APEC ($26 \pm 16 \mu\text{g m}^{-3}$ versus $39 \pm 23 \mu\text{g m}^{-3}$).
196 Sulfate, nitrate and ammonium before-APEC are 15 ± 13 , 28 ± 26 and $9.0 \pm 8.0 \mu\text{g m}^{-3}$ (Table 1)
197 and account for 8%, 16% and 5% of $PM_{2.5}$, respectively (Fig. 2). Their concentrations decrease
198 to 5.3 ± 2.8 , 10 ± 8.1 and $3.1 \pm 2.6 \mu\text{g m}^{-3}$ (Table 1) with the relative contributions to $PM_{2.5}$ down
199 to 5%, 10% and 3% during-APEC, respectively. While after-APEC their concentrations
200 increased to 11 ± 10 , 15 ± 13 and $6.9 \pm 6.4 \mu\text{g m}^{-3}$ and accounted for 7%, 9% and 4% of $PM_{2.5}$.
201 Such significant decreases in concentrations of OM and SIA during-APEC demonstrate the
202 efficiency of the emission controls. OC/EC ratio is almost constant during the whole period, but
203 WSOC/OC ratio decreased by 20% from 0.42 ± 0.13 before-APEC, 0.38 ± 0.16 during-APEC to
204 0.35 ± 0.17 after-APEC (Table 1). Since WSOC in fine aerosols consist mainly of secondary
205 organic aerosols (SOA) (Laskin et al., 2015), the decreasing ratio of WSOC/OC probably
206 indicates a reduced SOA production during the campaign.



207 **3.2 Oxalic acid and related SOA during the Beijing 2014 APEC campaign**

208 A homogeneous series of dicarboxylic acids (C_2 – C_{11}), keto-carboxylic acid and
209 α -dicarbonyls in the $PM_{2.5}$ samples were detected. As show in Table 2, total dicarboxylic acids
210 during the whole study period is $593 \pm 739 \text{ ng m}^{-3}$, which is lower than that observed during
211 Campaign of Air Quality Research in Beijing 2006 (CAREBeijing) (average 760 ng m^{-3}) and
212 2007 (average 1010 ng m^{-3}) (Ho et al, 2010, 2015) and the averaged wintertime concentration
213 reported by a previous research for 14 Chinese cities (904 ng m^{-3}) (Ho et al, 2007). Total
214 keto-carboxylic acid is $66 \pm 81 \text{ ng m}^{-3}$, while total dicarbonyls is $126 \pm 115 \text{ ng m}^{-3}$ (Table 2).
215 These values are higher than those during CAREBeijing 2006 and 2007 (Ho et al, 2010, 2015),
216 but close to the value observed for the 14 Chinese megacities (Ho et al, 2007). Being similar to
217 those previous observations, oxalic acid (C_2) is the most abundant diacid in the 2014 APEC
218 samples with an average of $334 \pm 461 \text{ ng m}^{-3}$ (range: 10 – 2127 ng m^{-3} , Table 2) during the whole
219 campaign, followed by methylglyoxal (mGly), succinin acid (C_4), terephthalic acid (tPh), and
220 glyoxal (Gly). These five species account for 43%, 10%, 9%, 6% and 6% of total detected
221 organic compounds (TDOC), respectively (Fig. 3).

222 As see in Fig. 4, TDOC in $PM_{2.5}$ are 1099 ± 1104 , 325 ± 220 and $487 \pm 387 \text{ ng m}^{-3}$ before-,
223 during- and after-APEC, respectively. In comparison with those before-APEC, TDOC
224 during-APEC decreased by 71%. Oxalic acid (C_2) is the leading species among the detected
225 organic compounds and accounted for 46%, 31% and 34% of TDOC during the three phases,
226 respectively (Fig. 4). C_2 is the end product of precursors that are photochemically oxidized in
227 aerosol aqueous phase via either oxidation of small compounds containing two carbon atoms or
228 decomposition of larger compounds containing three or more carbon atoms. Thus mass ratio of



229 C₂ to TDOC is indicative of aerosol aging (Wang et al., 2012; Ho et al., 2015). The highest
230 proportion of C₂ before- APEC suggests that organic aerosols during this period are more
231 oxidized. Glyoxal (Gly) and methylglyoxal (mGly) are the precursors of C₂. Mass ratios of both
232 compounds to TDOC are lowest before-APEC and highest during-APEC (Fig. 4). In comparison
233 with those before-APEC the decreased ratio C₂/TDOC and increased ratios of Gly/TDOC and
234 mGly/TDOC during- and after-APEC, together with a decreased ratio of WSOC/OC discussed
235 previously, further suggest a reducing production of SOA during the whole campaign.

236 3.3 Formation mechanism of oxalic acid

237 3.3.1 Correlation of oxalic acid with temperature, relative humidity (RH), aerosol liquid 238 water content (ALWC), acidity (pH), and sulfate

239 A few studies have pointed out that aerosol aqueous phase oxidation is a major formation
240 pathway for oxalic acid (Yu et al., 2005; Pinxteren, et al., 2014; Bikkina et al., 2015; Tilgner et al.,
241 2010). To explore the formation mechanism of oxalic acid, we calculated ALWC and pH of
242 PM_{2.5} aerosols by using ISOROPPIA-II, a state-of-the-art thermodynamic model (Weber et al.,
243 2016). As shown in Fig. 5, during the entire period C₂ showed a strong linear correlation with
244 sulfate ($R^2=0.70$ Fig. 5a), which is consistent with those observed in Xi'an (Wang et al., 2012)
245 and other Chinese cities (Yu et al., 2005). Previous studies on particle morphology showed that
246 sulfate particles internally mixes with SOA in Beijing especially in humid haze days (Li et al.,
247 2010, 2011), which probably indicates that they are formed via similar aqueous phase pathways.
248 Moreover, a robust correlation was also found for C₂ with RH ($R^2=0.64$, Fig. 5b) and aerosol
249 liquid water content (ALWC) ($R^2=0.61$, Fig. 5c), indicating that humid conditions are favorable
250 for the aqueous phase formation of C₂, which is most likely due to an enhanced gas-to-aerosol



251 aqueous phase partitioning of the precursors (e.g., Gly and mGly) (Fu et al., 2008; Wang et al.,
252 2015).

253 The H^+ aer calculated by using the ISORROPIA-II model indicates the aerosol aqueous
254 phase acidity. As seen in Fig. 5d, H^+ aer correlated well with SO_4^{2-} ($R^2 = 0.58$), probably
255 suggesting that SO_4^- is the key factor controlling the aerosol acidity. The linear fit slope (1.03)
256 indicates that the sulfate during the study period mostly appeared as bisulfate. In addition, H^+ aer
257 shows a significant positive correlation with C_2 ($R^2 = 0.84$) (Fig. 5e), possibly due to the fact that
258 acidic conditions are favorable for the formation of C_2 precursors. For example, Surratt et al
259 (2007; 2010) found that aerosol acidity can promote the formation of biogenic SOA (BSOA)
260 derived from isoprene oxidation such as 2-methylglyceric acid, Gly and mGly. These BSOA
261 precursors can be further oxidized into C_2 (Meng et al., 2014; Wang et al., 2009).

262 It is noteworthy that there is a significant positive correlation ($R^2 = 0.58$, $p < 0.001$) between
263 the mass ratios of C_2 /TDOC and ambient temperatures (Fig. 5f), which is similar to the results
264 found by previous researches (Ho et al., 2007; Strader et al., 1999) and suggests that aerosols are
265 more oxidized under higher temperature conditions (Erven et al, 2011; Carlton et al., 2009).
266 Therefore, both C_2 /TDOC ratios (31% versus 34%) and TDOC concentrations (278 ng m^{-3}
267 versus 426 ng m^{-3}) during- and after-APEC are lower than those before-APEC (Fig. 4) due to the
268 lower temperature conditions ($13 \pm 2.6 \text{ }^\circ\text{C}$, $7.0 \pm 1.7 \text{ }^\circ\text{C}$ and $4.3 \pm 1.3 \text{ }^\circ\text{C}$ before-, during- and
269 after-APEC periods, respectively) (Table1). Such results indicate that the significant reduction in
270 $PM_{2.5}$ during the APEC period is largely ascribed to the emission controls, but the favorable
271 meteorological conditions (e.g., the lower temperatures) may also take an important role.

272 3.3.2 Temporal variation in stable carbon isotopic composition of oxalic acid



273 To further discuss the formation mechanism of C_2 , we investigated the temporal variations
274 in concentration and stable carbon isotopic composition of C_2 in the $PM_{2.5}$ samples (Fig. 6).
275 Previous studies have demonstrated that Gly, mGly, glyoxylic acid (ωC_2) and pyruvic acid (Pyr)
276 are the precursors of C_2 (Carlton et al., 2006, 2007; Ervens and Barbara, 2004; Wang et al., 2012).
277 Thus, higher mass ratios of C_2 to its precursors indicate that organic aerosols are more oxidized
278 (Wang et al., 2010). As shown in Table 3, $\delta^{13}C$ of C_2 in this work positively correlated with the
279 mass ratios of $C_2/\omega C_2$, $C_2/mGly$ and TDOC/WSOC, demonstrating an enrichment of $\delta^{13}C$ during
280 the aerosol oxidation process. Because decomposition (or breakdown) of larger molecular weight
281 precursors in aerosol aqueous phase is the dominant formation pathway for C_2 in aerosol ageing
282 process (Kawamura et al., 2016; Gensch et al., 2014; Kirillova et al., 2013), during which
283 organic compounds release CO_2/CO by reaction with OH radical and other oxidants, resulting in
284 the evolved species enriched with lighter isotope (^{12}C) and the remaining substrate enriched in
285 ^{13}C due to kinetic isotope effects (KIEs) (Hoefs, 1997; Rudolph et al., 2002).

286 72-h backward trajectory analysis showed that air masses moved to Beijing during the
287 whole sampling period can roughly be categorized into three types (Fig. 6a) (all trajectories
288 during the entire study period can be found in the supplementary materials). (1) Regional type,
289 by which air masses originated from inland and east coastal China and moved slowly into
290 Beijing via a long distance within 72-h from its south regions, i.e., Henan, Shandong and Jiangsu
291 provinces. Air pollution in the three provinces is severe, and thus aerosols transported by this
292 type of air masses are of regional characteristics. (2) Local type, by which air masses originated
293 from Mongolia and north China, and moved quickly into Hebei province and then turned back to
294 Beijing. Air in Mongolia and north China is clean but is polluted in Hebei province, which is



295 adjacent to Beijing. Since the resident time of the air masses within Hebei province is very short,
296 thus aerosols transported by this type of air masses is of local characteristics and relatively fresh.

297 (3) Long-range transport type, by which air masses originated from Siberia and moved rapidly
298 into Beijing directly by a long-range transport. Aerosols from the long-range transport type air
299 masses are much more aged, while those from the local type air masses are fresh. Since severe
300 air pollution is widespread in the south regions, gas-to-aerosol phase partitioning of precursors
301 and subsequent aerosol-phase oxidation to produce SOA including C_2 continuously proceed
302 during the air mass movement. However, such a partition for producing SOA is not significant
303 when air mass move from Mongolia, north China and Siberia, because of the much less abundant
304 VOCs. In stead, aerosols in the north air masses are continuously oxidized, during which C_2 is
305 produced by photochemical decomposition of larger molecular weight precursors. Therefore, C_2
306 in $PM_{2.5}$ transported by the local type air masses are not only fresh and abundant but also
307 enriched in ^{12}C , whereas C_2 in $PM_{2.5}$ transported by the long-range transport type air masses are
308 aged, less abundant and enriched in ^{13}C due to KIE effects, as exemplified by the pink and light
309 blue columns in Fig. 6b, respectively. C_2 in $PM_{2.5}$ transported by the regional type air masses are
310 most abundant compared with that in other two type of air masses, which is not only due to the
311 severe air pollution in the Henan, Shandong and Jiangsu provinces but also due to the enhanced
312 photochemical oxidation under the humid and higher temperature conditions that occurred
313 mostly before-APEC, as discussed previously. Therefore, C_2 in the regional type air masses is
314 not only abundant but also enriched in ^{13}C (see black columns in Fig. 6b)

315 **3.4 Different chemical characteristics of $PM_{2.5}$ between two serve haze events**



316 From Fig. 1 and Table 4, it can be found that PM_{2.5} showed two equivalent maxima on 9th
317 October and 20th November during the whole study period. However, the chemical compositions
318 of PM_{2.5} during these two pollution events were significantly different. As shown in Fig. 7a,
319 relative abundances of SIA (sum of SO₄²⁻+NO₃⁻+NH₄⁺) to PM_{2.5} are 30% during the event I and
320 23% during the event II, respectively. The relative abundance of OM (21%, Fig. 7a) during the
321 event I is lower than that (37%) during the event II, but the ratios of WSOC/OC and TDOC/OC
322 are higher in the event I than in event II (Fig. 7b). Organic biomarkers in the PM_{2.5} have been
323 measured for the source apportionment (Wang et al., 2016) and cited here to further identify the
324 difference in chemical composition of PM_{2.5} between the two events. Levoglucosan is a key
325 tracer for biomass burning smoke. Mass ratio of levoglucosan to OC in PM_{2.5} (Lev/OC) is
326 comparable between the two events, suggesting a similar level of contributions of biomass
327 burning emission to PM_{2.5} before- and after-APEC. However, the mass ratios of PAHs and
328 hopanes to OC are lower in event I than those in event II (Fig. 7c), which again demonstrate the
329 enhanced emissions from coal burning for house heating, because these compounds are key
330 tracers of coal burning smokes (Wang et al., 2006). As seen in Fig. 7d, C₂ in the event I was
331 more enriched in δ¹³C. Such relatively more abundant SIA, WSOC and TDOC and heavier C₂ in
332 PM_{2.5} clearly demonstrate that PM_{2.5} during the event I were enriched with secondary products
333 while the fine particles during the event II were enriched with primary compounds. After-APEC
334 house heating activities including residential coal burning were activated, which emitted huge
335 amounts of SO₂, NO_x, and VOCs as well as primary particles. For example, Li et al (2015)
336 found VOCs at the urban center of Beijing rapidly increased from 48.28 ppbv during-APEC to
337 72.97 ppbv after-APEC. However, the lower temperature after-APEC is unfavorable for



338 photochemical oxidation, thus resulting in secondary aerosol production much less abundant
339 compared with those before-APEC. Temperatures during-APEC were also lower than those
340 before-APEC. Thus, compared to those before-APEC the much lower PM_{2.5} levels during-APEC
341 were caused not only by the emission controls but also by the lower temperatures.

342 **4. Summary and conclusion**

343 Temporal variations in molecular distribution of SIA, dicarboxylic acids, ketoacids and
344 α -dicarbonyl and stable carbon isotopic composition ($\delta^{13}\text{C}$) of C₂ in PM_{2.5} collected in Beijing
345 before-, during- and after- the 2014 APEC were investigated. Absolute concentrations and
346 relative abundances of SIA and C₂ in PM_{2.5} are highest before-APEC, followed by those after-
347 and during-APEC, suggesting that the fine aerosols before-APEC are enriched with secondary
348 products, mainly due to an enhanced photochemical oxidation under the higher temperature and
349 RH conditions. Concentrations of SIA, oxalic acid and related SOA in PM_{2.5} during-APEC are
350 2–4 times lower than those before-APEC, which is not only due to the effective emission
351 controls but also due to the lower temperature conditions that are unfavorable for secondary
352 aerosol production.

353 Positive correlations of C₂ with RH, sulfate mass, ALWC and aerosol acidity indicate that
354 C₂ formation pathway is involved an acid-catalyzed aerosol aqueous phase oxidation. SIA, C₂
355 and related SOA in the regional air masses are abundant with C₂ enriched in ¹³C. On the contrary,
356 those in the long-range transport air masses are much less abundant, although C₂ is also enriched
357 in ¹³C. By comparing the chemical composition of PM_{2.5} and $\delta^{13}\text{C}$ values of C₂ in two events that
358 are characterized by two highest values of PM_{2.5} before- and after-APEC, we further found that
359 compared with those before- APEC fine aerosols after-APEC are enriched with primary species



360 and C₂ is depleted in heavier ¹³C, although SO₂, NO_x and VOCs are much more abundant during
361 the heating season, again demonstrating the important role of temperature in secondary aerosol
362 production that is enhanced before-APEC and reduced during-APEC. Therefore, compared with
363 that before APEC the significant reduction of PM_{2.5} during-APEC is firstly ascribed to the
364 stringent emission controls and secondly attributed to the lower temperatures that suppressed
365 secondary aerosol production.

366

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375 References

- 376 Bikkina, S., Kawamura, K., and Miyazaki, Y.: Latitudinal distributions of atmospheric dicarboxylic acids,
377 oxocarboxylic acids, and α -dicarbonyls over the western North Pacific: Sources and formation pathways, *Journal*
378 *of Geophysical Research: Atmospheres*, 120, 5010-5035, 10.1002/2014jd022235, 2015.
- 379 Carlton, A. G., Turpin, B. J., Lim, H.-J., Altieri, K. E., and Seitzinger, S.: Link between isoprene and secondary
380 organic aerosol (SOA): Pyruvic acid oxidation yields low volatility organic acids in clouds, *Geophysical*
381 *Research Letters*, 33, 10.1029/2005gl025374, 2006.
- 382 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S., Reff, A., Lim, H.-J., and Ervens, B.: Atmospheric oxalic
383 acid and SOA production from glyoxal: Results of aqueous photooxidation experiments, *Atmos. Environ.*, 41,
384 7588-7602, 2007.
- 385 Carlton, A., Wiedinmyer, C., and Kroll, J.: A review of Secondary Organic Aerosol (SOA) formation from isoprene,
386 *Atmos. Chem. Phys.*, 9, 4987-5005, 2009.
- 387 Cheng, C., Wang, G., Meng, J., Wang, Q., Cao, J., Li, J., and Wang, J.: Size-resolved airborne particulate oxalic and
388 related secondary organic aerosol species in the urban atmosphere of Chengdu, China, *Atmospheric Research*,
389 161-162, 134-142, 10.1016/j.atmosres.2015.04.010, 2015.
- 390 Chow, J. C., Watson, J. G., Chen, L.-W. A., Chang, M. O., Robinson, N. F., Trimble, D., and Kohl, S.: The
391 IMPROVE_A temperature protocol for thermal/optical carbon analysis: maintaining consistency with a
392 long-term database, *Journal of the Air & Waste Management Association*, 57, 1014-1023, 2007.
- 393 Ervens, B.: A modeling study of aqueous production of dicarboxylic acids: 1. Chemical pathways and speciated
394 organic mass production, *J. Geophys. Res.*, 109, 10.1029/2003jd004387, 2004.
- 395 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and aqueous
396 particles (aqSOA): a review of laboratory, field and model studies, *Atmos. Chem. Phys.*, 11, 11069-11102, 2011.



- 397 Fu, T.-M., Jacob, D. J., Wittrock, F., Burrows, J. P., Vrekoussis, M., and Henze, D. K.: Global budgets of
398 atmospheric glyoxal and methylglyoxal, and implications for formation of secondary organic aerosols, *J.*
399 *Geophys. Res.*, 113, D15303, doi:10.1029/12007JD009505, 2008.
- 400 Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J., Yoo, K. Y., Beauchamp, J. L.,
401 Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle phase acidity and oligomer formation in secondary
402 organic aerosol, *Environ. Sci. Technol.*, 38, 6582-6589, 2004.
- 403 Gensch, I., Kiendler-Scharr, A., and Rudolph, J.: Isotope ratio studies of atmospheric organic compounds: Principles,
404 methods, applications and potential, *International Journal of Mass Spectrometry*, 365-366, 206-221,
405 10.1016/j.ijms.2014.02.004, 2014.
- 406 Guo, S., Hu, M., Zamora, M. L., Peng, J., Shang, D., Zheng, J., Du, Z., Wu, Z., Shao, M., Zeng, L., Molina, M. J.,
407 and Zhang, R.: Elucidating severe urban haze formation in China, *Proceedings of the National Academy of*
408 *Sciences*, 111(49), 17373-17378, 2014.
- 409 Ho, K. F., Cao, J. J., Lee, S. C., Kawamura, K., Zhang, R. J., Chow, J. C., and Watson, J. G.: Dicarboxylic acids,
410 ketocarboxylic acids, and dicarbonyls in the urban atmosphere of China, *J. Geophys. Res.*, 112,
411 10.1029/2006jd008011, 2007.
- 412 Ho, K. F., Lee, S. C., Ho, S. S. H., Kawamura, K., Tachibana, E., Cheng, Y., and Zhu, T.: Dicarboxylic acids,
413 ketocarboxylic acids, α -dicarbonyls, fatty acids, and benzoic acid in urban aerosols collected during the 2006
414 Campaign of Air Quality Research in Beijing (CAREBeijing-2006), *J. Geophys. Res.*, 115,
415 10.1029/2009jd013304, 2010.
- 416 Ho, K. F., Huang, R. J., Kawamura, K., Tachibana, E., Lee, S. C., Ho, S. S. H., Zhu, T., and Tian, L.: Dicarboxylic
417 acids, ketocarboxylic acids, α -dicarbonyls, fatty acids and benzoic acid in PM_{2.5} aerosol collected during
418 CAREBeijing-2007: an effect of traffic restriction on air quality, *Atmospheric Chemistry and Physics*, 15,
419 3111-3123, 10.5194/acp-15-3111-2015, 2015.
- 420 Hoefs, J., *Stable Isotope Geochemistry.*, Springer, New York, 1997.
- 421 Huang, R. J., Zhang, Y., Bozzetti, C., Ho, K. F., Cao, J. J., Han, Y., Daellenbach, K. R., Slowik, J. G., Platt, S. M.,
422 Canonaco, F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A.,
423 Schwikowski, M., Abbazade, G., Schnelle-Kreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El
424 Haddad, I., and Prevot, A. S.: High secondary aerosol contribution to particulate pollution during haze events in
425 China, *Nature*, 514, 218-222, 10.1038/nature13774, 2014.
- 426 Kawamura, K., and Watanabe, T.: Determination of stable carbon isotopic compositions of low molecular weight
427 dicarboxylic acids and ketocarboxylic acids in atmospheric aerosol and snow samples, *Analytical Chemistry*, 76,
428 5762-5768, 2004.
- 429 Kawamura, K., and Bikkina, S.: A review of dicarboxylic acids and related compounds in atmospheric aerosols:
430 Molecular distributions, sources and transformation, *Atmospheric Research*, 10.1016/j.atmosres.2015.11.018,
431 2015.
- 432 Kirillova, E. N., Andersson, A., Sheesley, R. J., Kruså, M., Praveen, P. S., Budhavant, K., Safai, P. D., Rao, P. S. P.,
433 and Gustafsson, Ö.: ¹³C- and ¹⁴C-based study of sources and atmospheric processing of water-soluble organic
434 carbon (WSOC) in South Asian aerosols, *J. Geophys. Res.*, 118, 614-626, 2013.
- 435 Li, J., Xie, S. D., Zeng, L. M., Li, L. Y., Li, Y. Q., and Wu, R. R.: Characterization of ambient volatile organic
436 compounds and their sources in Beijing, before, during, and after Asia-Pacific Economic Cooperation China
437 2014, *Atmos. Chem. Phys.*, 15, 7945-7959, 2015.
- 438 Li, W. J. and Shao, L. Y.: Mixing and water-soluble characteristics of particulate organic compounds in individual
439 urban aerosol particles. *J. Geophys. Res.* 115 (D02301), doi:10.1029/2009JD012575, 2010.



- 440 Li, W.J., Zhou, S.Z., Wang, X.F., Xu, Z., Yuan, C., Yu, Y.C., Zhang, Q.Z. and Wang, W.X.: Integrated evaluation
441 of aerosols from regional brown hazes over northern China in winter: Concentrations, sources, transformation,
442 and mixing states. *J. Geophys. Res.* 116 (D9), doi:10.1029/2010JD015099,2011.
- 443 Laskin, A., Laskin, J., and Nizkorodov, S. A.: Chemistry of atmospheric brown carbon, *Chem. Rev.*, 115,
444 4335-4382, 2015.
- 445 Meng, J., Wang, G., Li, J., Cheng, C., Ren, Y., Huang, Y., Cheng, Y., Cao, J., and Zhang, T.: Seasonal
446 characteristics of oxalic acid and related SOA in the free troposphere of Mt. Hua, central China: implications for
447 sources and formation mechanisms, *Sci Total Environ*, 493, 1088-1097, 2014.
- 448 Mu Quan, Z. S.: An evaluation of the economic loss due to the heavy haze during January 2013 in China., *China*
449 *Environmental Science*, 33, 2087-2094, 2013.
- 450 Pathak, R. K., Wang, T., Ho, K. F., and Lee, S. C.: Characteristics of summertime PM_{2.5} organic and elemental
451 carbon in four major Chinese cities: Implications of high acidity for water-soluble organic carbon (WSOC),
452 *Atmos. Environ.*, 45, 318-325, 10.1016/j.atmosenv.2010.10.021, 2011.
- 453 Rudolph, J., Czuba, E., Norman, A., Huang, L., and Ernst, D.: Stable carbon isotope composition of nonmethane
454 hydrocarbons in emissions from transportation related sources and atmospheric observations in an urban
455 atmosphere, *Atmos. Environ.*, 36, 1173-1181, 2002.
- 456 Strader, R., Lurmann, F., and Pandis, S. N.: Evaluation of secondary organic aerosol formation in winter, *Atmos.*
457 *Environ.*, 33, 4849-4863, 1999.
- 458 Sullivan, R. C., and Prather, K. A.: Investigations of the diurnal cycle and mixing state of oxalic acid in individual
459 particles in Asian aerosol outflow, *Environ. Sci. Technol.*, 41, 8062-8069, 2007.
- 460 Sun, Y., Wang, Z., Wild, O., Xu, W., Chen, C., Fu, P., Du, W., Zhou, L., Zhang, Q., Han, T., Wang, Q., Pan, X.,
461 Zheng, H., Li, J., Guo, X., Liu, J., and Worsnop, D. R.: "APEC Blue": Secondary Aerosol Reductions from
462 Emission Controls in Beijing, *Scientific reports*, 6, 20668, 10.1038/srep20668, 2016.
- 463 Surratt, J. D., Lewandowski, M., Offenberg, J. H., Jaoui, M., Kleindienst, T. E., Edney, E. O., and Seinfeld, J. H.:
464 Effect of Acidity on Secondary Organic Aerosol Formation from Isoprene, *Environ. Sci. Technol.*, 41,
465 5363-5369, 2007.
- 466 Surratt, J. D., Chan, A. W. H., Eddingsaas, N. C., Chan, M., Loza, C. L., Kwan, A. J., Hersey, S. P., Flagan, R. C.,
467 Wennberg, P. O., and Seinfeld, J. H.: Reactive intermediates revealed in secondary organic aerosol formation
468 from isoprene, *Proceedings of National Academy of Science of United States of America*, 107, 6640-6645,
469 2010.
- 470 Tang, G., Zhu, X., Hu, B., Xin, J., Wang, L., Munkel, C., Mao, G., and Wang, Y.: Impact of emission controls on
471 air quality in Beijing during APEC 2014: lidar ceilometer observations, *Atmos. Chem. Phys.*, 15, 12667-12680,
472 2015.
- 473 Tilgner, A. and Herrmann, H.: Radical-driven carbonyl-to-acid conversion and acid degradation in tropospheric
474 aqueous systems studied by CAPRAM. *Atmos. Environ.*, 44, 5415-5422, 2010.
- 475 Van Donkelaar, A., Martin, R. V., Brauer, M., Kahn, R., Levy, R., Verduzco, C., and Villeneuve, P. J.: Global
476 estimates of ambient fine particulate matter concentrations from satellite-based aerosol optical depth:
477 development and application, University of British Columbia, 2015.
- 478 van Pinxteren, D., Neusüß, C., and Herrmann, H.: On the abundance and source contributions of dicarboxylic acids
479 in size-resolved aerosol particles at continental sites in central Europe, *Atmos. Chem. Phys.*, 14, 3913-3928,
480 2014.
- 481 Wang, G., Cheng, C., Meng, J., Huang, Y., Li, J. and Ren, Y.: Field observation on secondary organic aerosols
482 during Asian dust storm periods: Formation mechanism of oxalic acid and related compounds on dust surface.
483 *Atmos. Environ.*, 113, 169-176, 2015.



- 484 Wang, G. H., Kawamura, K., Lee, S. C., Ho, K. F. and Cao, J. J.: Molecular, seasonal and spatial distributions of
485 organic aerosols from fourteen Chinese cities. *Environ. Sci. Technol.*, 40, 4619-4625, 2006.
- 486 Wang, G., Kawamura, K., Umemoto, N., Xie, M., Hu, S. and Wang, Z.: Water-soluble organic compounds in
487 PM_{2.5} and size-segregated aerosols over Mt. Tai in North China Plain. *Journal of Geophysical*
488 *Research-Atmospheres*, 114, D19208, doi:10.1029/2008JD011390, 2009.
- 489 Wang, G., Kawamura, K., Cheng, C., Li, J., Cao, J., Zhang, R., Zhang, T., Liu, S., and Zhao, Z.: Molecular
490 distribution and stable carbon isotopic composition of dicarboxylic acids, ketocarboxylic acids, and
491 alpha-dicarbonyls in size-resolved atmospheric particles from Xi'an City, China, *Environ. Sci. Technol.*, 46,
492 4783-4791, 2012.
- 493 Wang, G., Xie, M., Hu, S., Gao, S., Tachibana, E., and Kawamura, K.: Dicarboxylic acids, metals and isotopic
494 compositions of C and N in atmospheric aerosols from inland China: implications for dust and coal burning
495 emission and secondary aerosol formation, *Atmos. Chem. Phys.*, 10, 6087-6096, 2010.
- 496 Wang, G., Wang, J., Ren, Y. and Li, J.: Chemical characterization of organic aerosols from Beijing during the 2014
497 APEC (under preparation), 2016.
- 498 Wang, Y., Zhang, Q. Q., He, K., Zhang, Q., and Chai, L.: Sulfate-nitrate-ammonium aerosols over China: response
499 to 2000–2015 emission changes of sulfur dioxide, nitrogen oxides, and ammonia, *Atmos. Chem. Phys.*, 13,
500 2635-2652, 2013.
- 501 Weber, R. J., Guo, H., Russell, A. G., and Nenes, A.: High aerosol acidity despite declining atmospheric sulfate
502 concentrations over the past 15 years, *Nature Geoscience*, 9, 282-285, 10.1038/ngeo2665, 2016.
- 503 Xing, L., Fu, T. M., Cao, J. J., Lee, S. C., Wang, G. H., Ho, K. F., Cheng, M. C., You, C. F., and Wang, T. J.:
504 Seasonal and spatial variability of the OM/OC mass ratios and high regional correlation between oxalic acid and
505 zinc in Chinese urban organic aerosols, *Atmos. Chem. Phys.*, 13, 4307-4318, 2013.
- 506 Xu, W. Q., Sun, Y. L., Chen, C., Du, W., Han, T. T., Q. Wang, Q., Fu, P. Q., F. Wang, Z., Zhao, X. J., Zhou, L. B., Ji,
507 D. S., Wang, P. C., and Worsnop, D. R.: Aerosol composition, oxidation properties, and sources in Beijing:
508 results from the 2014 Asia-Pacific Economic Cooperation summit study, *Atmos. Chem. Phys.*, 15, 13681–13698,
509 2015.
- 510 Yu, J. Z., Huang, X.-F., Xu, J., and Hu, M.: When Aerosol Sulfate Goes Up, So Does Oxalate: Implication for the
511 Formation Mechanisms of Oxalate, *Environ. Sci. Technol.*, 39, 128-133, 2005.
- 512 Zhang, Q., Jimenez, J. L., Worsnop, D. R., and Canagaratna, M.: A case study of urban particle acidity and its
513 influence on secondary organic aerosol, *Environ. Sci. Technol.*, 41, 3213-3219, 2007.
- 514 Zhang, Q., Streets, D. G., Carmichael, G. R., He, K., Huo, H., Kannari, A., Klimont, Z., Park, I., Reddy, S., and Fu,
515 J.: Asian emissions in 2006 for the NASA INTEX-B mission, *Atmos. Chem. Phys.*, 9, 5131-5153, 2009.
- 516 Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.: Formation of
517 urban fine particulate matter, *Chemical Reviews*, 115, 3803-3855, 2015.
- 518 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D.,
519 Pöschl, U., Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic
520 weather, regional transport and heterogeneous reactions, *Atmos. Chem. Phys.*, 15, 2969-2983, 2015.
- 521

522 **Figure Captions**

523

524 **Figure 1.** Temporal variations of meteorological conditions, gaseous pollutants and major
525 components of $PM_{2.5}$ during the 2014 APEC campaign. (The brown shadows
526 represent two air pollution events characterized by highest $PM_{2.5}$ levels before- and
527 after-APEC, while the blue shadow represents the APEC event).

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529 **Figure 2.** Chemical composition of $PM_{2.5}$ during the 2014 APEC campaign.

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531 **Figure 3.** Molecular distributions of dicarboxylic acids and related compounds in $PM_{2.5}$ of
532 Beijing, China during the 2014 APEC campaign. The pie chart is the average
533 composition of total detected organic compounds (TDOC) and the top number is the
534 average mass concentration of TDOC of the whole study period.

535

536 **Figure 4.** Compositions of total detected organic compounds (TDOC) in $PM_{2.5}$ during the 2014
537 APEC campaign.

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539 **Figure 5.** Correlation analysis for oxalic acid (C_2) and sulfate in $PM_{2.5}$ during the whole 2014
540 APEC campaign. **(a-c)** Concentrations of C_2 with sulfate, relative humidity (RH), and
541 aerosol liquid water content (ALWC); **(d, e)** sulfate and C_2 with aerosol acidity [H^+]
542 and **(f)** temperature with mass ratio of C_2 to total detected organic compounds
543 (C_2 /TDOC).

544

545 **Figure 6.** **(a)** 72-h backward trajectories determined by the National Oceanic and Atmospheric
546 Administration Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPLIT)
547 model arriving at the sampling site to reveal the major air mass flow types during the
548 study period. Northwesterly wind (light blue) was most frequently (64%), followed by
549 northerly (21%, pink) and southerly (15%, black) and is defined as long-range
550 transport, local and regional type, respectively (see the definitions in the text); **(b)** Time
551 series of $\delta^{13}C$ values and concentration of oxalic acid during the whole study period
552 (Colors in Fig. 6a are corresponding to those in Fig. 6b).

553

554 **Figure 7.** Comparison of chemical composition of $PM_{2.5}$ during two air pollution events. **(a)**
555 Percentages of major species in $PM_{2.5}$; **(b, c)** mass ratios of major species and organic
556 tracers in $PM_{2.5}$; **(d)** stable carbon isotope composition of oxalic acid (C_2) (Data about
557 levoglucosan (Lev), PAHs and hopanes are cited from Wang et al (2016)).

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 Table 1. Meteorological parameters and concentrations ($\mu\text{g m}^{-3}$) of gaseous pollutants and chemical components of $\text{PM}_{2.5}$ in Beijing during the 2014 APEC campaign

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	Whole period (N=48)	Before-APEC (08/10–02/11) (N=26)	During-APE (03/11–12/11) (N=10)	After-APEC (13/11–14/11) (N=12)
I. Meteorological parameters				
Temperature, °C	9.5±4.3 (3.0–18)	13±2.6 (9.0–18)	7.0±1.7 (4.0–10)	4.3±1.3 (3.0–7.0)
Relative humidity, %	56±19 (17–88)	62±19 (22–88)	47±14 (17–65)	51±16 (29–80)
Visibility, km	8.8±6.8 (1.0–28)	7.3±6.6 (1.0–24)	13±7.7 (6.0–28)	7.2±4.2 (2.0–15)
Wind speed, km h ⁻¹	8.0±4.9 (3.0–26)	7.6±4.8 (3.0–26)	9.4±6.6 (3.0–26)	7.8±2.9 (3.0–13)
II. Gaseous pollutants				
O ₃	48 ± 23 (6.0–115)	55 ± 24 (9.0–115)	52 ± 13 (25–69)	29 ± 18 (6.0–60)
SO ₂	12 ± 8.5 (2.0–43)	8.8 ± 4.6 (2.0–19)	7.6 ± 3.9 (2.0–15)	23 ± 8.8 (13–43)
NO ₂	68 ± 29 (10–135)	71 ± 27 (22–118)	45 ± 18 (10–69)	78 ± 29 (45–135)
III. Major components of PM _{2.5}				
PM _{2.5}	157 ± 110 (16–408)	178 ± 122 (16–408)	98 ± 46 (28–183)	161 ± 100 (36–383)
SO ₄ ²⁻	12 ± 11.5 (1.2–43)	15 ± 13 (1.2–43)	5.3 ± 2.8 (1.8–11)	11 ± 10 (2.9–34)
NO ₃ ⁻	21 ± 22 (0.32–88)	28 ± 26 (0.32–88)	10 ± 8.1 (1.2–26)	15 ± 13 (2.9–46)
NH ₄ ⁺	7.3 ± 7.2 (0.2–28)	9.0 ± 8.0 (0.21–28)	3.1 ± 2.6 (0.2–8.6)	6.9 ± 6.4 (1.0–22)
OC	28 ± 18 (5.7–78)	26 ± 16 (6.0–67)	19 ± 7.6 (5.7–29)	39 ± 23 (9.7–78)
EC	8.8 ± 5.4 (1.4–25)	8.6 ± 4.6 (1.4–18)	6.0 ± 2.7 (1.5–9.6)	12 ± 7.0 (2.1–25)
WSOC	10 ± 6 (2.4–32)	11 ± 4.6 (3.1–32)	6.4 ± 2.6 (2.4–11)	11 ± 6.1 (4.5–24)
NO ₃ ⁻ /SO ₄ ²⁻	1.6±0.8 (0.3–4.3)	1.7±0.9 (0.3–4.3)	1.6±0.7 (0.5–2.4)	1.4±0.4 (0.8–2.2)
OC/EC	3.3±0.6 (2.2–4.7)	3.2±0.7 (2.2–4.5)	3.3±0.6 (2.0–4.3)	3.4±0.5 (2.7–4.7)
WSOC/OC	0.39±0.15 (0.10–0.71)	0.42±0.13 (0.13–0.71)	0.38±0.16 (0.16–0.65)	0.35±0.17 (0.10–0.63)



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Table 2. Concentrations of dicarboxylic acids and related compounds in PM_{2.5} in Beijing during the 2014 APEC campaign (ng m⁻³)

	Whole period (N=48)	Before-APEC (08/10–02/11) (N=26)	During-APEC (03/11–12/11) (N=10)	After-APEC (13/11–14/11) (N=12)
I. Dicarboxylic acids				
Oxalic, C ₂	334 ± 461 (10–2127)	502 ± 564 (10.5–2127)	101 ± 69 (35–251)	166 ± 157 (22–554)
Malonic, C ₃	31 ± 42 (ND–247)	45.7 ± 52.1 (1.44–247)	12 ± 8.0 (3.4–22.8)	16 ± 10.9 (ND–36)
Succinic, C ₄	74 ± 118 (3.0–722)	111 ± 150 (3.0–722)	24 ± 14 (7.1–42)	36 ± 26 (4.9–90)
Glutaric, C ₅	8.7 ± 12 (ND–68)	13 ± 15 (ND–68.1)	2.9 ± 2.24 (0.9–5.8)	4.9 ± 4.2 (ND–13)
Adipic, C ₆	13 ± 14 (0.9–83)	17 ± 18 (1.9–83)	5.9 ± 3.8 (2.1–14)	9.9 ± 7.1 (2.0–23)
Pimelic, C ₇	2.1 ± 3.8 (ND–27)	2.6 ± 5.1 (ND–27)	1.1 ± 0.7 (0.2–2.3)	2.0 ± 1.1 (0.9–4.4)
Suberic, C ₈	10 ± 11 (ND–66)	12 ± 13 (ND–66)	7.6 ± 5.0 (1.3–16)	8.7 ± 6.0 (2.0–21)
Azelaic, C ₉	5.0 ± 4.9 (0.5–21)	6.4 ± 5.7 (0.6–21)	1.7 ± 0.9 (0.5–3.2)	4.6 ± 3.3 (1.3–13)
Sebacic, C ₁₀	7.7 ± 7.4 (ND–34)	9.4 ± 8.8 (ND–34)	4.2 ± 3.6 (0.5–11)	6.8 ± 4.9 (1.4–16)
Undecanedioic, C ₁₁	11 ± 13 (ND–77)	14 ± 16 (ND–77)	3.3 ± 2.5 (ND–7.5)	9.4 ± 6.4 (0.8–23)
Methylsuccinic, iC ₅	13 ± 16 (0.6–79)	18 ± 19 (0.6–79)	4.8 ± 3.0 (1.0–9.2)	8.4 ± 6.0 (2.3–19)
Methylglutaric, iC ₆	7.5 ± 10 (ND–36)	11 ± 12 (ND–36)	0.9 ± 0.9 (ND–2.6)	4.6 ± 5.1 (ND–14)
Maleic, M	3.4 ± 3.9 (ND–15)	4.6 ± 4.7 (ND–15)	1.4 ± 0.8 (ND–2.9)	2.4 ± 2.0 (ND–6.3)
Fumaric, F	7.2 ± 8.8 (ND–64)	10 ± 11 (ND–64)	2.2 ± 1.5 (ND–5.4)	4.7 ± 3.2 (1.4–10)
Phthalic, Ph	17 ± 14 (1.5–64)	20 ± 16 (1.5–64)	10 ± 6.8 (2.3–20)	17 ± 9.0 (6.4–31)
Isophthalic, iPh	2.1 ± 2.5 (ND–10)	2.9 ± 2.8 (ND–10)	2.0 ± 2.1 (0.2–5.9)	0.5 ± 0.3 (ND–3.2)
Terephthalic, tPh	46 ± 35 (2.6–133)	50 ± 35 (2.6–123)	28 ± 19 (4.7–59)	53 ± 40 (7.4–133)
Subtotal	593 ± 739 (25–3788)	849 ± 905 (25–3788)	214 ± 135 (72–447)	354 ± 279 (85–965)
II. Keto-carboxylic acids				
Pyruvic, Pyr	24 ± 20 (1.3–84)	31 ± 23 (2.4–84)	15 ± 12 (1.3–36)	15 ± 9.3 (3.2–33)
Glyoxylic, ωC ₂	33 ± 51 (1.2–300)	48 ± 64 (1.2–300)	10 ± 7.7 (2.6–21)	20 ± 23 (2.8–80)
7-Oxoheptanoic, ωC ₇	8.8 ± 14 (ND–90)	13 ± 17 (ND–90)	4.2 ± 3.6 (ND–13)	4.5 ± 5.1 (ND–17)
Subtotal	66 ± 81 (3.6–474)	92 ± 99 (3.6–474)	30 ± 22 (5.9–66)	40 ± 35 (13–128)
III α-Dicarbonyls				
Glyoxal, Gly	44 ± 47 (4.2–270)	57 ± 56 (4.2–270)	22 ± 19 (4.9–47)	35 ± 30 (7.3–101)
Methylglyoxal, mGly	82 ± 82 (ND–406)	102 ± 96 (ND–406)	60 ± 52 (15–139)	58 ± 51 (5.8–144)
Subtotal	126 ± 115 (5.3–466)	158 ± 132 (5.3–466)	81.6 ± 67.4 (22–186)	93 ± 80 (14–225)
TDOC ^b	785 ± 872 (36–4636)	1099 ± 1104 (36–4636)	325 ± 220 (107–664)	487 ± 387 (117–1318)

^aND: not detectable; ^bTDOC: total detected organic compounds.



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Table 3 Linear correlation coefficients of $\delta^{13}\text{C}$ of C_2 with $\text{C}_2/\omega\text{C}_2$, C_2/mGly , and TDOC/WSOC

	$\text{C}_2/\omega\text{C}_2$	C_2/mGly	TDOC/WSOC
$\delta^{13}\text{C}$	0.49**	0.35*	0.41*

** $p < 0.01$; * $p < 0.05$



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Table 4. Meteorological parameters and chemical compositions ($\mu\text{g m}^{-3}$) of two maximum $\text{PM}_{2.5}$ between two pollution episodes in Beijing

	T ($^{\circ}\text{C}$)	RH (%)	V ^a (km)	$\text{PM}_{2.5}$	OC	EC	SIA ^b	TDOC ^c
Event I (8/10-11/10, Before-APEC)	16.7 ± 0.8	82 ± 4	1.5 ± 0.5	349 ± 57	45 ± 12	12 ± 2	106 ± 39	2749 ± 1357
Event II (18/11-21/11, After-APEC)	4.5 ± 1.7	62 ± 13	3.5 ± 1.5	259 ± 102	60 ± 21	17 ± 6	60 ± 32	831 ± 400

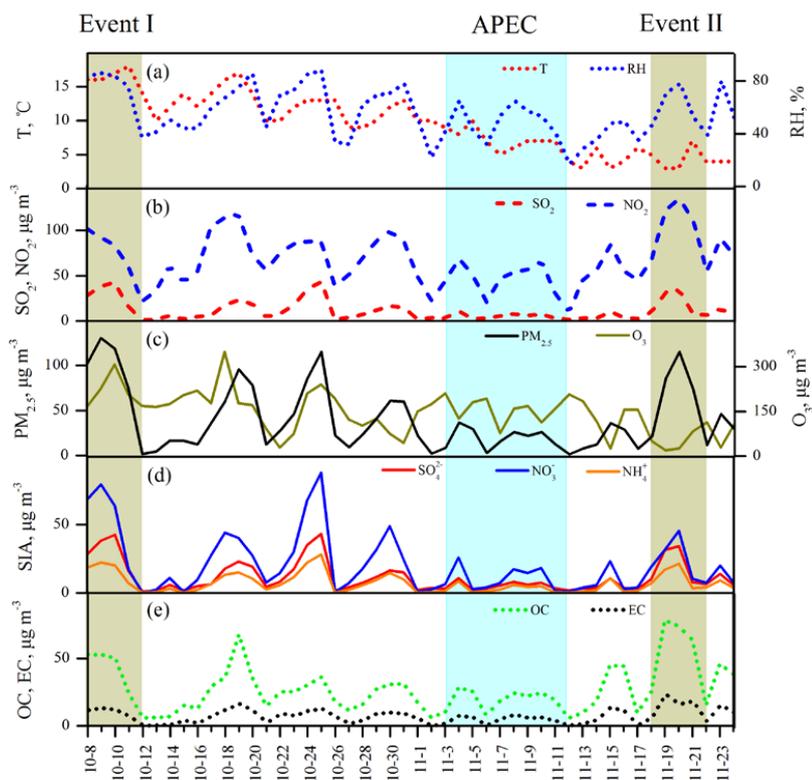
^aV: visibility; ^bSIA: secondary inorganic ions (the sum of sulfate, nitrate and ammonium); ^cTDOC: total detected organic compounds

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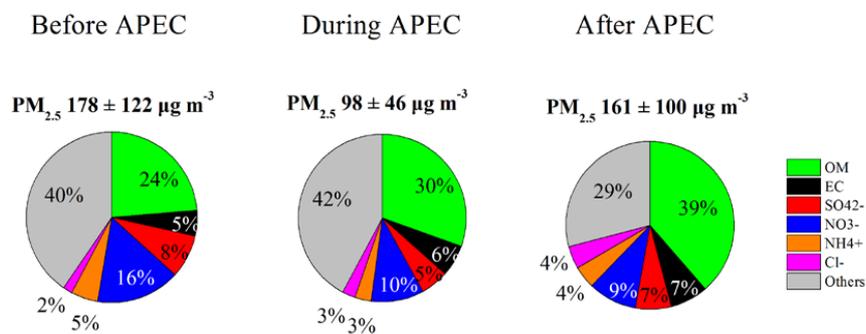


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596 Figure 1. Temporal variations of meteorological conditions, gaseous pollutants and major components of $PM_{2.5}$
597 during the 2014 APEC campaign. (The brown shadows represent two air pollution events characterized
598 by highest $PM_{2.5}$ levels before- and after-APEC, while the blue shadow represents the APEC event).



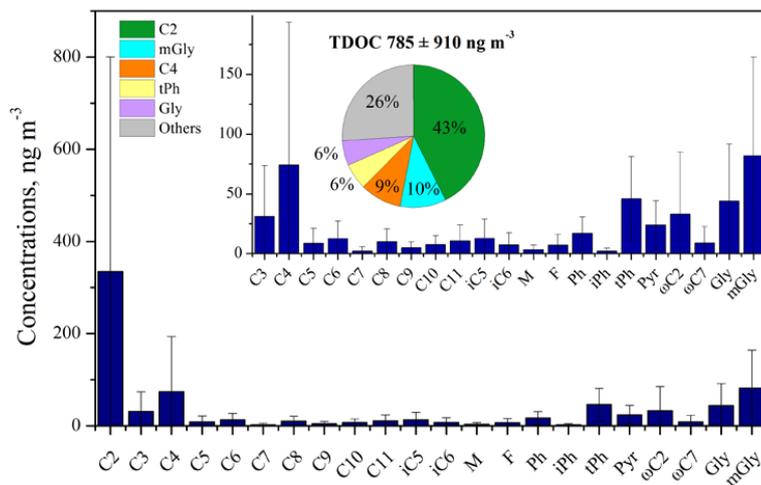
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600 **Figure 2.** Chemical composition of PM_{2.5} during the 2014 APEC campaign.

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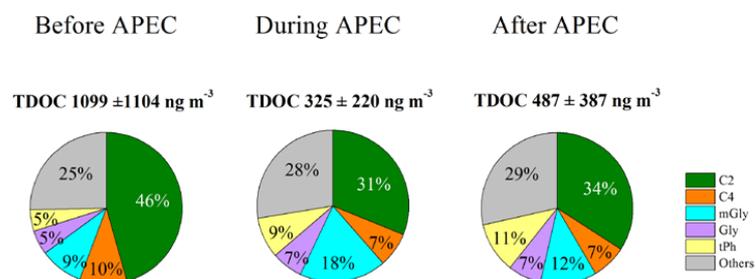


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604 **Figure 3.** Molecular distributions of dicarboxylic acids and related compounds in $\text{PM}_{2.5}$ of Beijing, China during the
605 2014 APEC campaign. The pie chart is the average composition of total detected organic compounds
606 (TDOC) and the top number is the average mass concentration of TDOC of the whole study period.
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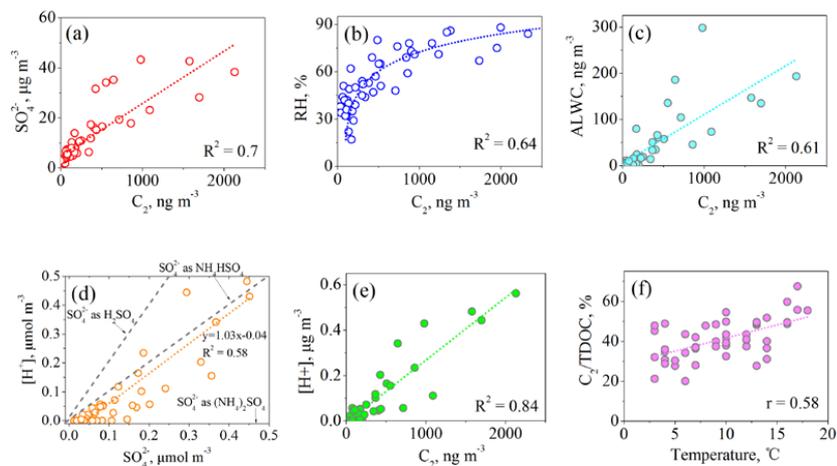
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Figure 4. Compositions of total detected organic compounds (TDOC) in PM_{2.5} during the 2014 APEC campaign.



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614 **Figure 5.** Correlation analysis for oxalic acid (C_2) and sulfate in $PM_{2.5}$ during the whole 2014 APEC campaign. (a-c)

615 Concentrations of C_2 with sulfate, relative humidity (RH), and aerosol liquid water content (ALWC); (d, e)

616 sulfate and C_2 with aerosol acidity $[H^+]$ and (f) temperature with mass ratio of C_2 to total detected organic

617 compounds (C_2 /TDOC).

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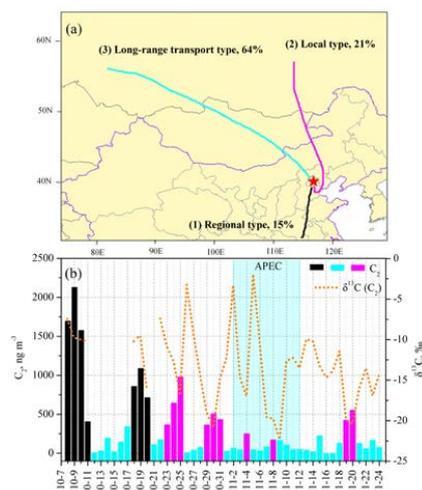
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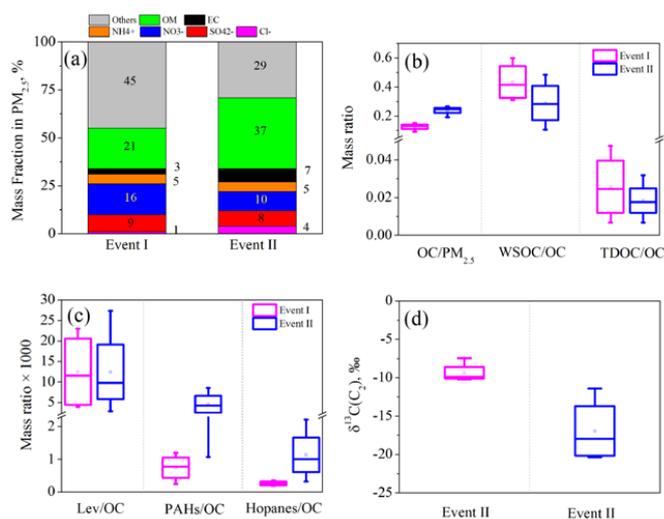
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Figure 6. (a) 72-h backward trajectories determined by the National Oceanic and Atmospheric Administration Hybrid Single Particle Lagrangian Integrated Trajectory (HYSPPLIT) model arriving at the sampling site to reveal the major air mass flow types during the study period. Northwestern wind (light blue) was most frequently (64%), followed by northerly (21%, pink) and southerly (15%, black) and is defined as long-range transport, local and regional type, respectively (see the definitions in the text); (b) Time series of $\delta^{13}\text{C}$ values and concentration of oxalic acid during the whole study period (Colors in Fig. 6a are corresponding to those in Fig. 6b).



638

639 **Figure 7.** Comparison of chemical composition of PM_{2.5} during two air pollution events. (a) Percentages of major
 640 species in PM_{2.5}; (b, c) mass ratios of major species and organic tracers in PM_{2.5}; (d) stable carbon
 641 isotope composition of oxalic acid (C₂) (Data about levoglucosan (Lev), PAHs and hopanes are cited
 642 from Wang et al (2016)).

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