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The impact of Pacific Decadal Oscillation on springtime dust activity in Syria

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Abstract. The increasing trend of aerosol optical depth in the Middle East and a recent severe dust storm in Syria have raised questions as whether dust storms will increase and promoted investigations on the dust activities driven by the natural climate variability underlying the ongoing human perturbations such as the Syrian civil war. This study examined the influences of the Pacific decadal oscillation (PDO) on dust activities in Syria using an innovative dust optical depth (DOD) dataset derived from Moderate Resolution Imaging Spectroradiometer (MODIS) Deep Blue aerosol products. A significantly negative correlation is found between the Syrian DOD and the PDO in spring from 2003-2015. High DOD in spring is associated with lower geopotential height over the Middle East, Europe, and North Africa, accompanied by near surface anomalous westerly winds over the Mediterranean basin and southerly winds over the eastern Arabian Peninsula. These large-scale patterns promote the formation of the cyclones over the Middle East to trigger dust storms and also facilitate the transport of dust from North Africa, Iraq, and Saudi Arabian to Syria, where the transported dust dominates the seasonal mean DOD in spring. A negative PDO not only creates circulation anomalies favorable to high DOD in Syria but also suppresses precipitation in dust source regions over the eastern and southern Arabian Peninsula and northeastern Africa.

On the daily scale, in addition to the favorable large-scale condition associated with a negative PDO, enhanced atmospheric instability in Syria associated with increased precipitation in Turkey and northern Syria is also critical for the development of strong springtime dust storms in Syria.

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1. Introduction

Dust aerosol is an important component in the climate system that can modify global and regional energy and water balances (e.g., Tegen et al., 1996; Miller and Tegen, 1998; Miller et al., 2004; Lau et al., 2009; Yue et al., 2010; 2011; Choobari et al., 2014; Huang et al., 2014). Dust particles interact with both solar and terrestrial radiation, modifying temperature profile and hydrological cycle, which impact regional and global climate. For instance, studies found that mineral dust influences the strength of the West African monsoon (e.g., Miller and Tegen, 1998; Miller et al., 2004; Yoshioka et al., 2007; Solmon et al., 2008; 2012; Mohowald et al., 2010; Strong et al., 2015) and Indian summer monsoonal rainfall (Vinoj et al., 2014; Jin et al., 2014, 2015, 2016; Solmon et al., 2015; Kim et al., 2016). Dust particles can also serve as ice cloud nuclei and influence the microphysical and macrophysical properties of the cloud (e.g., Levin et al., 1996; Rosenfeld et al., 1997; Wurzler et al., 2000; Nakajima et al., 2001; Bangert et al., 2012), including its droplet size, number concentration, lifetime, albedo, and in turn affecting the regional radiative budget and hydrological cycle. Mineral dust also provides nutrients for ocean phytoplankton, affecting ocean productivity and therefore carbon and nitrogen cycles and ocean albedo (e.g., Fung et al., 2000; Shao et al., 2011). Strong dust storms also have severe social and health impacts (e.g., Morman and Plumlee, 2013), affecting public transportation and causing damage to the eye and lung.

The Middle East is one of the dustiest regions in the world, and recent study suggested an increasing trend of aerosol optical depth largely due to dust emission (e.g., Pozzer et al., 2015; Klingmüller et al., 2016). A once-in-ten-years severe dust storm recently occurred in Syria during September 6-9th, 2015, and raised many attentions. More than one thousand people in

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Syria were hospitalized due to breathing difficulties*. The causes of such a strong dust storm are not fully understood, but there have been speculations that the ongoing civil war in Syria is responsible for it. The argument is that crop fields were abandoned or destroyed by the war, so soil dust is easier to be uplifted by wind from these unprotected land fields. More importantly, will severe dust storms like the one above increase in the future?

Notaro et al. (2015) studied the dust activities over the Arabian Peninsula and related the increased dust activities during 2007-2013 to the persistent dry condition over the "Fertile Crescent" (namely Syria, Iraq, Israel, and Jordan) primarily caused by a combined effect of La Niña and a negative phase of Pacific decadal oscillation (PDO). Associated with the drought is crop failure and increased dust activities over the Arabian Peninsula. While others, such as Chin et al. (2014) found that the positive trend of dust emission over the Middle East from 2000-2009 is related to increased surface wind speed, and Klingmüller et al. (2016) attributed the positive trend of aerosol optical depth (AOD) over the region (mainly Saudi Arabia, Iraq, and Iran) from 2001-2012 to a combined effects of decreased precipitation and soil moisture over Iraq and adjacent areas and enhanced surface wind over the Africa Red Sea coastal area that increased dust emission.

The above studies suggested that remote sea surface temperature (SST), local precipitation, surface wind, and vegetation all influence the dust activities over the Middle East. Among these factors, the remote forcings, such as tropical Pacific SST, the PDO, not only affect the precipitation variations over the region but can also influence local circulation including near surface winds, both of which influence dust emission. It is thus quite important to understand the

* https://www.washingtonpost.com/news/worldviews/wp/2015/10/07/syrias-war-helped-create-an-epic-dust-storm-scientists-say/

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influences of these low-frequency long-lasting forcings on dust activities underlying the ongoing human perturbations such as civil war, land use change, and anthropogenic emission.

In this paper we examine the influence of the PDO on the variations of dust activities in Syria from 2003-2015 using Moderate Resolution Imaging Spectroradiometer (MODIS) Deep Blue dust optical depth (DOD). Previous studies on the connection between the PDO and dust activities in the Arabian Peninsula mainly focus on precipitation (e.g., Notaro et al., 2015; Yu et al., 2015). Here we explore thoroughly how the PDO influences the key factors associated with the dust activities in Syria on the interannual and daily scales.

The following section presents the data and methodology used in the paper. The co-variations of DOD in Syria and PDO are presented in sections 3 and their physical connections are analyzed in details in section 4. In section 5, we examined the key factors associated with strong spring dust events in Syria using daily reanalysis and observations. To what extend can climate models capture the connections between the PDO and Syrian dust activities is examined using the Geophysical Fluid Dynamics Laboratory (GFDL) AM3 model and discussed in section 6. Major conclusions are summarized in section 7.

2. Data and methodology

2.1 Satellite and observational datasets

2.1.1 Dust optical depth

Daily and monthly dust optical depth data is derived from MODIS aerosol products retrieved using Deep Blue (M-DB2) algorithm, which employs radiance from the blue channels to detect aerosols over bright land surface (e.g., desert). Because surface reflectance is low at blue channels, increases of reflectance and spectral contrast indicate the presence of aerosols

(Hsu et al., 2004, 2006). Ginoux et al. (2012) used collection 5.1 level 2 aerosol products from MODIS aboard Aqua satellite to derive DOD. Here, both MODIS aerosol products (collection 6, level 2) from the Aqua and Terra platforms are used for cross-validation. Terra passes the equator from north to south around 10:30 a.m. local time while Aqua passes the equator from south to north around 13:30 p.m. local time. While the frequency of the maximum daily 10 m wind in Syria peaks in the afternoon, later than the passing time of both Terra and Aqua, the averaged maximum total wind speed is nearly evenly distributed over a day (supplementary Figure S1). Thus, the results from both platforms are likely to provide a more complete picture of dust activities than either one alone.

Aerosol products such as AOD, single scattering albedo, Angström exponent are first interpolated to a regular 0.1° by 0.1° grid using the algorithm described by Ginoux et al. (2010). The dust optical depth is then derived from AOD following the methods of Ginoux et al. (2012) with adaptions for the newly released MODIS Collection 6 aerosol products. To separate dust from other aerosols, we use the Angström exponent, which has been shown to be highly sensitive to particle size (Eck et al., 1999) and single scattering albedo which is less than one for dust due to its absorption of solar radiation. Instead of using negative values of Angström exponent as by Ginoux et al. (2012), we use a continuous function relating Angström exponent to fine-mode aerosol optical depth established by Anderson et al. (2005; their Eq. 5) based on ground-based data. The DOD data is available from January 2003 to December 2015.

Daily and monthly DOD indices are formed by averaging DOD data in Syria between 34°-36.5°N and 36.5°-41°E to characterize dust activities. The averaging area covers most region of Syria. We also tested using a smaller averaging box (33.5° -36°N, 36.5°-39°E) for the DOD index, and the results are similar.

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2.1.2 Precipitation

Version 7 of Tropical Rainfall Measurement Mission (TRMM) Multi-satellite Precipitation Analysis (TMPA) daily product (3B42) is used. This product covers from 50°S to 50°N with a spatial resolution of 0.25° by 0.25° and is available from 1998 to present. Several important changes are applied to version 7 product, including using additional satellite data such as early record of Microwave Humidity Sounder (MHS) and operational Special Sensor Microwave Imager (SSM/I) record, using a new infrared brightness temperature dataset before the start of the Climate Prediction Center (CPC) 4-km Merged Global IR data set, using a single uniformly processed surface precipitation gauge analysis, using a latitude-band calibration scheme for all satellites, and adding output fields in the data files (Huffman and Bolvin, 2014).

Precipitation Reconstruction over Land (Chen et al., 2002; hereafter PRECL) from National Oceanic and Atmospheric Admiration (NOAA) is a global analysis available monthly from 1948 to present at a 1° by 1° resolution. The dataset is derived from gauge observations from the Global Historical Climatology Network (GHCN), version 2, and the Climate Anomaly Monitoring System (CAMS) datasets.

The monthly precipitation of the Climatic Research Unit (CRU) time-series (TS) 3.23 (Harris et al. 2014) is also used as a supplement to the PRECL. CRU TS 3.23 dataset covers 1901-2014, with a spatial resolution of 0.5° by 0.5° over land (excluding Antarctica). The gridded data is based on the analysis of over 4,000 individual weather station records.

2.1.3 Temperature

The Hadley Centre sea ice and sea surface temperature (HadISST) data set (Rayner et al., 2003) from the UK Met Office is available monthly from 1870 to the present with a horizontal resolution of 1° by 1° grid. Monthly SST from HadISST and land surface temperature from CRU TS 3.23 (0.5° by 0.5°) from 1948-2015 are used to examine temperature patterns associated with dust activities.

2.1.4 Leaf area index (LAI)

LAI characterizes the canopies of plants. It is defined as the one-sided green leaf area per unit ground area in broadleaf canopies and as half the total needle surface area per unit ground area in coniferous canopies. LAI at zero is considered as bare ground while around 10, dense forests. Monthly LAI is derived from the version 4 of Climate Data Record (CDR) of Advanced Very High Resolution Radiometer (AVHRR) surface reflectance (Claverie et al., 2014) and produced by the National Aeronautics and Space Administration (NASA) Goddard Space Flight Center (GSFC) and the University of Maryland. The gridded monthly data is on a 0.05° by 0.05° horizontal resolution and available from 1981 to present. A detailed discussion on the algorithm and evaluation of the dataset can be found by Calverie et al. (2016).

Monthly MODIS LAI level 4 data on the Aqua platform (MYD15A2) is also used for 2003-2015. The original data files were obtained via personal communication (Ranga Myneni and Taejin Park; Boston University) and then reprocessed to fill the missing data by Paul Ginoux. The horizontal resolution of the data is 0.1° by 0.1° degrees.

2.2 Reanalyses

Daily and monthly geopotential height, horizontal winds, specific humidity from the National Centers for Environmental Prediction (NCEP)/National Center for Atmospheric Research (NCAR) reanalysis (Kalnay et al., 1996, hereafter NCEP1) from 1948 to 2015 are used.

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Its horizontal resolution is 2.5° by 2.5° and have 17 vertical levels from 1000 hPa to 10 hPa, with 8 levels between 1000 hPa and 300 hPa. This reanalysis is used primarily in this study due to its long record.

ERA-interim (Dee et al., 2011) from the European Centre for Medium-Range Weather Forecasts (ECMWF) is a global reanalysis with a horizontal resolution of T255 (about 0.7° or 80 km) and 37 vertical levels, available from 1979 to present. Monthly and four (two) times daily analysis (forecast) variables are used. The time coverage of ERA-Interim is shorter than the NCEP1 but its high resolutions supplements the latter. Monthly horizontal winds and geopotential heights are compared with the NCEP1 in the same period (1979-2015) and show similar features (see discussion in section 3). Daily 10 meter and 850 hPa winds, and forecast variables such as precipitation are used to investigate the key factors associated with dust activities.

2.3 Climate indices

The PDO and Niño3.4 indices are downloaded from the website of NOAA Climate Prediction Center (http://www.esrl.noaa.gov/psd/data/climateindices/list/). The monthly standardized PDO index is derived from the leading principal component of SST anomalies in the northern Pacific Ocean (20°N north). The monthly mean of global mean SST anomalies are removed in the PDO index, thus the influence of global warming is not included. The data is available from 1948 to present. The monthly Niño 3.4 index is derived from the extended reconstructed sea surface temperature (ERSST v4; Huang et al., 2015; Lui et al., 2015; Huang et al., 2016) averaged over the tropical Pacific between 5°N–5°S and 120°–170°W and is available from 1950 to present.

2.4 Model output

To examine the relationship between the Syrian dust activities and the PDO, the output from the atmospheric component (AM3) of a general circulation model (CM3; Donner et al., 2011) developed at the GFDL is used. The finite-volume algorithms described in Lin and Rood (1996, 1997) and Lin (1997, 2004) is used in the dynamic core in AM3. Different from earlier versions of the model, a general curvilinear coordinate system is used and largely improved the computational efficiency. A hybrid vertical coordinate (Simmons and Burridge, 1981) of 48 layers is used, with the top model layer at about 1 Pa (~86 km). AM3 calculates the mass distribution and optical properties of aerosols according to their emission, chemical production, transport, and dry and wet deposition. The dust source function follows the scheme of Ginoux et al. (2001), which places preferential sources in topographic depressions. The simulated aerosol optical depth and co-albedo show improved correlations with the AErosol RObotic NETwork (AERONET) station observations than earlier version of the model (CM2.1) but slightly underestimate AOD in the Middle East.

A historical run is conducted using the observed monthly SSTs from the Hadley center to drive the AM3. The simulated wind is nudged toward the NCEP/NCAR reanalysis with a relaxation timescale of 6 hours (Moorthi and Suarez, 1992), similar to the method used by Li et al. (2008). A moderate horizontal resolution of 2° by 2.5° is used. The simulation was conducted from 1948 to 2010. Results from 1960 to 2010 are presented.

3. The co-variations between springtime DOD in Syria and the PDO during 2003-2015

Notaro et al. (2015) related low-frequency variations of monthly precipitation over Syria and Iraq to the El Niño–Southern Oscillation (ENSO) and the PDO and thus built the connection

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between dust activities and Pacific SSTs. Here we focus on Syria. Figure 1 shows the time series of monthly Syrian DOD indices (seasonal cycle removed) from both MODIS Aqua (green) and Terra (grey) platforms and the negative PDO index from January 2003 to December 2015. The variations of the DOD and the negative PDO indices are quite similar, showing strong decadal variations underlying interannual variations. Both are relatively weak during 2003-2007, relatively strong during 2007-2012 and become relatively weak since 2013. Note that the DOD indices increase again in 2015 in association with the severe dust storm in September. The correlation between monthly DOD indices and PDO index are -0.51 (Aqua) and -0.50 (Terra) from 2003-2015. Other indices, such as Niño 3.4 and Syrian LAI also have significant but lower correlations with the DOD indices (Table 1).

ENSO is known to influence precipitation over the Middle East, by decreasing precipitation in La Niña years and increasing precipitation in El Niño years (e.g., Price et al., 1998; Mariotti et al., 2005; 2007, Chakaborty et al., 2005; Barlow et al., 2002, Wang et al., 2014; Yu et al., 2015; Banerjee and Kumar, 2016), and its influence is generally stronger in La Niña years (e.g., Wang et al., 2014). Previous study also showed a comparable influence of the PDO on precipitation over the "Fertile Crescent" region including Syria and Iraq versus that by ENSO (correlations of 0.52 versus -0.57 from 1979-2013, Notaro et al., 2015). The correlations in this study indicate that the PDO plays a greater role than ENSO in modulating dust activities in Syria in the recent decade. The PDO is not completely independent of ENSO, but can be viewed as a phenomenon driven by multiple physical processes, including the tropical Pacific SST, atmospheric noise, Aleutian low, Kuroshio-Oyashio Extension, Pacific-North American pattern, Rossby wave breaking and etc. (e.g., Evans et al., 2001; Newman et al., 2003; Schneider and Cornuelle, 2005; Strong and Magnusdottir, 2009; Mills and Walsh, 2013). While some modeling

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studies suggested that up to half of the variance of the PDO can be explained by ENSO (e.g., Alexander et al., 2002; Liu and Alexander, 2007), others found that there certain part of extra-tropical Pacific SST variability are totally independent of ENSO (e.g., Zhang et al., 1997; Deser and Blackmon, 1995; Zhang and Delworth, 2015). Here the correlation between the monthly PDO and Niño 3.4 indices is 0.63 from January 2003 to December 2015, suggesting that statistically the Niño 3.4 index explains about 40% of the variances of the PDO index.

Figure 2 shows the correlation between the PDO index and Aqua and Terra DODs in spring, along with correlations between the Syrian DOD indices and PDO index in individual month and on annual mean from 2003-2015. The correlation pattern between the PDO index and Aqua DOD is very similar to that associated with Terra DOD (Figs. 2a-b), with negative correlations over most areas of Syria and a stronger correlation over the eastern Syria than the western part. DOD over Iraq and northern Saudi Arabia is also significantly negatively correlated with the PDO. The correlation between the monthly PDO and Syrian DOD indices shows a persistent negative relationship through the year, with higher correlations during March-April-May and also in July-August and December (Fig. 2c). The low correlation in September is due to the severe dust storm in 2015. The seasonal mean correlations between the DOD indices and the PDO index in MAM are -0.90 for both the Aqua DOD and Terra DOD indices, again much higher than their correlations with other indices (Table 2).

The connection between the Syrian DOD and PDO is further examined in Figure 3, which shows the correlations between the Syrian DOD indices and SST and land surface temperatures during MAM and on the annual mean, along with SST and land surface temperature patterns associated with the PDO index. As shown in Figs. 3a-b, the SST pattern associated with the positive phase of the PDO in spring are quite similar to that in the annual

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mean, with anomalous warm SST over the tropical Pacific and along the east basin of North Pacific and anomalous cold SST over the subtropical central to western North Pacific. SST in the India Ocean is also positively associated with the PDO. But the correlation between land surface temperature and the PDO is not significant in most regions, including the Arabian Peninsula.

Correlations between Aqua DOD index and SST in MAM are nearly opposite to those of the PDO, with negative correlations over the tropical and eastern Pacific but positive correlations over the central North Pacific (Fig. 3c). Correlations between the annual mean Aqua DOD index and SST and land surface temperature are similar to that in spring, but with slightly weaker magnitude over the central Pacific (Fig. 3d). The temperature patterns correlated with the Terra DOD index are quite similar to that of Aqua DOD index (Figs. 3e-f). Both DOD indices are not significantly correlated with the land surface temperature over the Middle East in spring.

Since Figs. 2-3 indicate a stronger negative correlation between the PDO and Syrian DOD index during MAM than on annual mean or other seasons, the following analysis focuses on their connections in spring. Spring is also the time when Syrian DOD is high and dust storms are most active (supplementary Figures S2-3).

4. How does the PDO affect Syrian dust activities in spring?

In this section we explore the mechanisms underlying the strong negative correlations between the PDO and Syrian DOD indices. Since MODIS DOD only covers the recent decade and the PDO has low-frequency decadal variations with cycles of about 15-25 and 50-70 years (e.g., Minobe, 1997; Mantua and Hare, 2002), it is difficult to assess their long-term (i.e., beyond the recent decade) relationship directly. To establish their connections, we first compare the patterns of circulation and precipitation associated with DOD and PDO indices during 2003-

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2015. The patterns associated with long-term variations of the PDO (e.g., 1948-2015 from the NCEP1 or 1979-2015 from the ERA-interim) are then examined and compared with those associated with the PDO in the recent decade. The PDO shows a dominant negative phase during the late 1940s to 1970s and turns into a dominant positive phase during the 1980s and 1990s and becomes mostly negative again since the middle 2000s (e.g., JISAO/University of Washington website[†]). Therefore the NCEP1 reanalysis covers about one 50-70 years multi-decadal cycle of the PDO, while the EAR-Interim covers about half of the cycle. Assuming that the factors dominated the variations of Syrian DOD persist beyond the recent decade (e.g., also valid during 1948-2015), then the similarity between the patterns associated with the PDO in the recent decade and those over longer time periods would indicate a persistent influence of the PDO on dust activities in Syria.

How does the PDO affect the circulation over the Middle East? Figure 4 shows the regression of 200 hPa (contours) and 850 hPa (shading) geopotential heights onto the standardized PDO and DOD indices. Fig. 4a shows in the positive phase of PDO, stationary wave propagates from the North Pacific through North America, northern Atlantic and Europe to the Middle East. The regressions at 850 hPa are generally similar to those at 200 hPa, indicating an equivalent barotropic structure. Anomalous positive geopotential height over the Arabian Peninsula is associated with a positive PDO during 2003-2015 (Fig. 4a) and also during 1948-2015 (Fig. 4c). The geopotential height patterns associated with the both Aqua and Terra DOD indices are nearly opposite to those associated with the positive PDO index, with anomalous high over the central North Pacific and northeastern North America and anomalous low over the west coast of North America and Europe and the Middle East (Figs. 4b, d).

[†] http://research.jisao.washington.edu/pdo/graphics.html

Fig. 4 suggests that the PDO can influence the variations of springtime DOD in Syria through its modification on the circulation over the Middle East. A negative PDO reduces the geopotential height both at 850 and 200 hPa over the Arabian Peninsula, a scenario that favors high DOD in spring.

The connection between Syrian DOD and 850 hPa winds and geopotential heights are further examined in Figure 5. Fig. 5a shows that a positive PDO is associated with anomalous easterly winds north of 40°N, anomalous northerly winds over the central Mediterranean Sea around 15°E and weak southeasterly winds over the eastern Mediterranean Sea and western Syria in the recent decade (Fig. 5a) and from 1948-2015 (Fig. 5c). Anticyclonic winds are located over Oman and Yemen at the south coast of Arabian Peninsula accompanied by an anomalous high. Positive geopotential height anomalies are also located over the northwestern Africa, and East Africa. On the other hand, winds associated with high Syrian DOD indices are anomalous westerly over the northern Mediterranean basin and Syria (Figs. 5b and d). Anomalous cyclonic flows are located over the southern Arabian Peninsula, nearly opposite to that associated with the positive phase of the PDO. The overall lower geopotential height over the Middle East and Africa also facilitate the formation of the cyclones (such as Sharav cyclones) that are important for the spring peak of dust storms in the northern Arabian Peninsula (e.g., Shao et al. 2001; Israelevich, et al. 2003; Dayan et al. 2008; Dayan et al. 2012).

The regression patterns of 850 hPa winds and geopotential height onto the PDO and DOD indices in the ERA-Interim are generally consistent with those shown in the NCEP1 (Supplementary Figure S4).

Next we examine the associated variations of near surface wind and precipitation that are tied to these geopotential height and low-level wind patterns. Figure 6 shows the regression of

10 m winds (vectors) and the cubic 10 m wind speed (shading) in the ERA-Interim onto standardized PDO index and onto the Aqua and Terra DOD indices. The ERA-Interim is chosen here because of its higher horizontal resolution compared to the NCEP1, and thus is more suitable to examine surface wind variations associated with dust blasting in small scales. Cubic wind speed is used here as classical dust emission scheme relates dust flux to the third power of 10 m horizontal winds. The patterns of the surface wind associated with the DOD indices (Figs. 6b and d) are largely similar to that of winds at 850 hPa (Figs. 5 and S4). Anomalous southwesterly winds from coastal North Africa and over the Mediterranean Sea tend to bring dust from North Africa to Syria. Such a route of dust transport has not been directly examined, but was discussed in back trajectories studies on the airflow patterns onto Israel (e.g., Dayan, 1986). Earlier studies also have suggested a transport of dust from North Africa to the Mediterranean basin (e.g., D'Almeida, 1986; Moulin et al., 1998; Kubilay et al., 2000). Anomalous northerly flow over the Red Sea and the west coast of the Arabian Peninsula (and a weaker wind speed) and anomalous southerly flow (and a stronger wind speed) over the eastern Peninsula also suggest a transport of dust from the source regions in the middle and southern Arabian Peninsula, e.g., An Nafud and Ad Dahna deserts, dry riverbeds such as Al Batin, Al-Rimah, and Al Sahba and Rub's al Khali Sandy desert in Saudi Arabia (Ginoux et al., 2012). An Nafud and Ad Dahna deserts are major sources of dust storms in Saudi Arabian in spring to early summer (Notaro et al., 2013) and can also be an important source for Syrian DOD. A modeling study on the sources of spring DOD in Syria also confirms that North Africa (including Libya, Algeria, and Egypt) is the largest source with the secondary source from the Arabian Peninsula (Iraq and Saudi Arabia), and the overall transported DOD is much higher than local DOD in Syria in spring (Ginoux et al., in preparation).

Variations of surface wind associated with the PDO are different from that associated with high DOD, with nearly opposite pattern of cubic wind speeds over the Arabian Peninsula (Figs. 6a and c). When the PDO is positive, anomalous northerly winds pass through the Mediterranean Sea and turn into westerly over the northeastern Africa, which may bring dust from Africa to Israel, Jordan, and Syria as well as increase the moisture transport from the Mediterranean Sea (Fig. 6a). The anomalous southerly wind from the Red Sea also brings moisture onto Syria and tends to promote precipitation. Over southwestern Syria, the anomalous southerly wind is from the less dusty area over northwestern Saudi Arabia and is less likely to enhance Syrian DOD. A weak anomalously westerly flow over the northeastern Saudi Arabia around 30°N and 45°E tends to block the northward transport of dust from the southern and middle Arabian Peninsula to Syria. This westerly flow weakens if extending the time period of regression to 1979 (Fig. 6c). A stronger anomalously westerly flow along the west coast of Saudi Arabia and a southerly flow from the south coast of the Arabian Peninsula are also found in association with the positive PDO during 1979-2015, bringing moisture onto Syria. The discrepancies between the regression patterns in the recent decade and those during 1979-2015, e.g., over the eastern Mediterranean Sea (Figs. 6a and c), are likely associated with the decadal variations of the PDO, indicating an instable connection between the PDO and surface winds in some areas.

The anomalous precipitation pattern correlated with the PDO and DOD indices are shown in Figure 7. The PDO has a mild connection with precipitation over the Middle East in spring on the interannual time scale. Similar correlation patterns are found in previous studies that correlated the PDO with low-frequency (i.e., >7years) variations of precipitation (e.g., Dai, 2013, Fig. 2c). A positive phase of the PDO is associated with increased precipitation over most area of the Arabian Peninsula, Turkey, western Iran, and northeastern Africa over Libya and Egypt in the recent decade and during 1948-2015 (Figs. 7a and c). Patterns are similar during 1979-2015 (not shown). On the other hand, high DOD in Syria is associated with reduced precipitation over the Arabian Peninsula (particularly, Iraq, central and southern Saudi Arabia), Egypt, eastern Algeria, and western Libya. This is consistent with Fig. 6, which indicates dust transport from these areas to Syria in the spring.

Figure 8 shows the vertically integrated mass-weighted moisture flux (vector) and its magnitude (shading) onto the standardized PDO and DOD indices in MAM from 2003-2015 and also onto the PDO index during 1948-2015. Consistent with the precipitation regression patterns (Fig. 7), anomalous anticyclonic moisture flux brings more moisture onto the Arabian Peninsula from the Red Sea and Gulf of Aden associated with a positive PDO (Figs. 8a and c). Moisture flux over the northeastern Africa is also enhanced. Oppositely, high DOD in Syria is associated with an anomalous cyclonic flux centered over the south coast of the Arabian Peninsula, which reduces moisture flux onto the southern Arabian Peninsula (Figs. 8b and d). Moisture flux over northeastern Africa is also reduced associated with an anomalous cyclonic flow over Sudan and Egypt. The pattern of the anomalous moisture fluxes associated with PDO and DOD indices are quite similar to that of the 850 hPa winds (e.g., Fig 5), indicating a dominant role played by low-level moisture transport.

These anomalous circulation and moisture flux patterns are also linked to the stability of lower atmosphere. Figure 9 shows the regression of low-level Moist Static Energy (MSE; integrated from surface to 700 hPa) onto the standardized PDO and DOD indices. MSE is defined as a sum of sensible, latent, and geopotential energy in a column air, i.e., $MSE = c_pT + Lq + gz$, where c_p is the specific heat of air at constant pressure, T is air temperature, L is the

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latent heat of vaporization of water, q is specific humidity, and g is the gravity acceleration, and z geopotential height. MSE increasing with altitude denotes a stable atmosphere, so high MSE in the lower atmosphere indicates an instable condition, and vice versa. The patterns of the anomalous MSE are tied to the changes of moisture flux and precipitation anomalies. Reduced MSE is located over large areas of the Arabian Peninsula, particularly, the southwest coast, in association with high DOD in Syria, indicating a more stable low-level atmosphere and thus less precipitation (Fig. 9b and d). Such a low MSE is also found over North Africa, but in a weaker magnitude. The pattern associated with a positive PDO is nearly opposite, with increased low-level MSE over the Arabian Peninsula, Red Sea, and along the east coast Egypt, denoting an instable atmosphere associated with anomalous moisture transport (Figs. 8a and c) and promoting convection and precipitation (Figs. 9a and c).

In short, Figs. 4-9 show spring circulation and precipitation patterns favorable to high DOD in Syria. Anomalous low pressure over Europe, the southern Arabian Peninsula, and northeastern to eastern Africa promotes westerly winds from North Africa and southerly flow over the southeastern Arabian Peninsula, both of which tend to transport dust to Syria. The anomalous moisture fluxes associated with the geopotential height and wind anomalies also favor a dry and stable condition over the dust source regions adjacent to Syria, such as Saudi Arabia, Iraq, and North Africa. The circulation and precipitation patterns associated with a positive PDO are largely opposite to those associated with high DOD in Syria in the recent decade, which explains the strong negative correlation between the two. Examination on circulation and precipitation variations associated with the PDO in a longer time period (either from 1979-2015 or 1948-2015) show generally similar patterns, but also with some discrepancies. If the conditions associated with high DOD in Syria are valid beyond the recent decade, i.e., 2003-2015, the negative role of the PDO on spring dust activities in Syria is also likely to persist.

5. Analysis on strong dust storms in spring

Severe dust storms usually only persist a few days or even a few hours (e.g., Haboobs) and seasonal or monthly averages reduce the variability of dust activities and may smooth out some important features. In this section we discuss the conditions associated with strong dust storms in Syria using daily DOD and reanalysis variables and compare these conditions with those from seasonal mean patterns discussed above including the teleconnections with the PDO. Composites are formed based on daily Syrian DOD index from Aqua. Days during March, April, and May from 2003-2015 are selected when the DOD index is greater than 1 standard deviation to form daily composites. Results are very similar but patterns are in slightly stronger (weaker) magnitudes if choosing the threshold of 1.5 (0.5) standard deviations (not shown).

Figure 10 shows the composite of Aqua and Terra daily DOD (shading) and ERA-Interim 10 m winds based on the Aqua DOD index. The patterns are quite similar in Aqua and Terra DODs. DOD anomaly is above 0.3 over Syria and western Iraq (Figs. 10a-b). Anomalous high DOD is also located over eastern Saudi Arabia and the northeastern Africa, indicating a possible transport of dust from these areas to Syria. Anomalous cyclonic flow is centered over southern Turkey around 30°E, with anomalous strong westerly wind blowing from North Africa and southerly flow from the eastern Arabian Peninsula, consistent with dust transport discussed in the above section. However, different from the seasonal mean regression patterns (e.g., Fig. 6), there are the anomalous southerly winds from the Red Sea and Persian Gulf.

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Figure 11 shows the composite of daily precipitation from ERA-Interim two times daily forecast and TRMM daily precipitation for days with DOD anomaly above 1 standard deviation. Precipitation anomalies are shown in percentages (with reference to the climatological mean) instead of absolute values in order to highlight the precipitation variations over the Middle East, where the magnitude of precipitation in spring is quite low (less than 1 mm day^{-1} in most of the areas). Patterns of precipitation anomalies associated with strong dust storms are quite similar in the ERA-Interim and TRMM, but with greater magnitudes in the ERA-Interim. Precipitation increases significantly (about 80% in TRMM and more than 100% in the ERA-Interim) over Turkey and the northeastern Mediterranean, but decreases over the central and southern Arabian Peninsula and northeastern Africa. Syria sits in between the anomalous wet and dry regions, with slightly increased precipitation in its northern domain. These features are somewhat similar to the results of previous studies on strong dust storms and our understanding on the seasonal mean patterns associated with high DOD in Syria. Strong dust storms such as Haboobs (usually about 1 kilometer height and tens to hundreds of kilometer length) are usually associated with convective storms (e.g., Miller et al., 2008; Roberts and Knippertz, 2012; Vukovic et al., 2014; Dempsey, 2014). The cold downdraughts from convective storms spread out and can lift the dust from the surface to form a dusty towering "wall" as the front of a Haboob. Similarly, severe precipitation and convection in Turkey and northern Syria can produce an unstable atmospheric condition in the region, and the intensified low-level winds can lift dust from the surface and thus increase DOD. Reduced precipitation over southern Arabian Peninsula and North Africa facilitates dust transport from these source areas to Syria.

Figure 12 shows composites for vertically integrated mass weighted moisture flux (vectors) and its magnitude (shading) from the NCEP1 and convective available potential energy

(CAPE) from the ERA-Interim for days with Aqua DOD index greater than 1 standard deviation. Fig. 12a shows anomalous westerly flux from northern Egypt and southerly flux from the Red Sea and Persian Gulf largely increase the moisture transport to Syria and eastern Turkey, while the reduced moisture fluxes along the south coast of Iran, southern Arabian Peninsula, southern Red Sea, and the Gulf of Aden are quite similar to those patterns association with high Syrian DOD in spring (Figs. 8b and d) and a negative PDO (i.e., opposite to Fig. 8c). Consistently, CAPE is increased over Turkey and Syria, indicating an unstable atmospheric condition associated with increased moisture transport to the region, while over southern Arabian Peninsula and northeastern Africa where moisture flux is reduced CAPE is decreased (Fig. 12b).

The connection between the PDO and daily strong dust storms is also verified by correlating an index of the occurrence of strong dusty events (i.e., daily DOD anomaly greater than 1 standard deviation) in MAM with HadISST from 2003-2015 (supplementary Figure S5).

Figs. 10-12 suggest that severe dust storms occur under both favorable large-scale and regional-scale features. Remote forcing such as PDO modifies springtime circulation and precipitation patterns, e.g., a negative phase of PDO decreases precipitation over the southern Arabian Peninsula, northeastern Africa and favors the transport of dry dusty air from these regions to Syria, while strong convective storms over Turkey favor the dust lifting and formation of strong dust storms in Syria.

6. Can the recent GFDL climate model capture the connection between the PDO and Syrian DOD?

To what extent can current climate model capture the connection between the PDO and DOD? We examined such relationships in the GFDL AM3 model. Figure 13a shows the

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correlation between modeled DOD index and surface temperature from 1960-2010 in MAM. The correlation pattern over the North Pacific is quite similar to that of a negative PDO (e.g., Fig. 3c), but only significant over the northern North Pacific, indicting a weaker such connection in the model. Over land, DOD is highly positively associated with surface temperature in the northern to northeastern Africa, which is not seen in the observations, and may suggest an overestimation of the connection in the model. The correlation pattern is similar if calculated from 1951 to 2010, but with a slightly stronger positive correlation over the tropical eastern Pacific (not shown).

Fig. 13b shows the regression of standardized DOD index onto 200 hPa (contours) and 850 hPa geopotential heights (shading) during MAM 1960-2010. The wave trains propagation from north Pacific to the Middle East is quite similar to that shown in the NCEP1 using observed DOD index (Figs. 4b and d) but in a weaker magnitude in the tropical and subtropical North Pacific, consistent with the weak SST correlations (Fig. 13a). The anomalous low over Africa dips down to the eastern Sahel, probably in association with the biased correlation between the DOD and surface temperature over the northeastern Africa.

Figure 13 suggests that AM3 can partially capture the connection between the dust activities in Syria and the PDO. A few reasons may contribute to this underestimation. First, AOD is slightly underestimated in the Middle East compared with AERONET, and the simulated DOD is also less than that in MODIS, which indicates that dust variability may be underestimated in the region. Current dust scheme in AM3 only relates dust emission to dust source map and surface wind speed, while the influence of soil moisture on dust emission is not explicitly considered. Thus, the anomalous dust transport from southern Arabian Peninsula and northeast Africa in association with the dry conditions under a negative PDO may not be captured by the model. Our analysis also suggest that the DOD in the model is highly correlated with dry deposition over the western Mediterranean, along the north coast of Egypt, and Turkey and with the southwesterly winds in these region, indicating an very strong connection with the dust sources in Africa, which may be an overestimation. A new dust emission scheme that considers the influences of soil moisture and vegetation cover and land use change is currently under development, and the relationship between Syrian DOD and PDO is likely to be better represented in this newer version of the model.

7. Conclusions

Dust activities in the Middle East have been related to a lot factors, such as remote sea surface temperatures, near surface winds, and precipitation variability. The ongoing civil war and a recent severe dust storm in Syria in 2015 raised concerns as whether dust activities will increase in the region. First step toward answering this question is to understand the dust activities driven by the natural climate variability. Here we examine the connection between Syrian dust activities and the Pacific decadal oscillation (PDO) using innovative dust optical depth (DOD) datasets retrieved from MODIS Deep Blue aerosol products and multiple observations and reanalyses.

A significantly negative correlation is found between Syrian DOD and the PDO in springtime during 2003-2015, suggesting that the PDO index explains about 81% variances of Syrian DOD in spring in the recent decade. Such a connection is revealed not only on precipitation as emphasized by previous studies (e.g., Yu et al., 2014; Notaro et al., 2015) but also on other aspects such as the circulation patterns and surface winds. It is found that high DOD in Syria during spring is associated with low geopotential height over Europe, southern Arabian Peninsula, and northeastern to eastern Africa. Associated with these anomalous height patterns are the westerly wind anomalies over the Mediterranean basin and southerly wind anomalies over the southeastern Arabian Peninsula, favoring dust transport from these regions to Syria, where the transported dust dominates the DOD in spring (Ginoux et al., in preparation). A positive PDO is connected with wind and height patterns largely opposite to those associated with high DOD over Syria. Positive phase of the PDO also tends to increase precipitation over the Arabian Peninsula and northeastern Africa via anomalous moisture transport that increases moisture supply and also reduces the stability of low-level atmosphere. A negative PDO thus is not only associated with wind and geopotential height patterns favorable to high DOD in Syria but also tends to reduces precipitation in the dust source regions such as Iraq, Saudi Arabia, and northeastern Africa, and thus favors dust transport to Syria. This explains why the correlation between the Syrian DOD index and the PDO index is much higher than other individual index such as precipitation, leaf area index, and 10 m winds in Syria (Tables 1-2). The influences of the PDO on circulation and precipitation patterns over the Middle East largely persist beyond the recent decade, i.e., over 1948-2015, but also show some exceptions. The lack of long-term observations also brings uncertainties to the connection between the PDO and Syrian DOD.

Different from the patterns on seasonal mean discussed above, analysis on the daily composites of strong spring dust storms shows both the influence of the PDO and local features. In spring, strong dust storms (DOD anomalies greater than 1 standard deviation) in Syria is associated with an anomalous cyclonic flow centered over the northeastern Mediterranean Sea and Turkey, and southerly wind anomalies from the Red Sea and Persian Gulf. Consistently, moisture flux onto Turkey and Syria is enhanced and thus destabilizes the atmosphere and promotes precipitation in Turkey and convection and dust uplifting in Syria. Meanwhile, reduced moisture fluxes onto the southern Arabian Peninsula and east coast of Egypt and Sudan in

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association with a negative PDO favor a dry and stable condition in Saudi Arabia and northeastern Africa, facilitating dust transport from these regions to Syria.

We examined the teleconnection between the DOD and PDO in the GFDL AM3 model. A weaker connection compared to that in the observation is found, which may be partially related to model's underestimation of the mean DOD and its variability in this area. The new dust scheme that includes the influence of soil moisture and precipitation is likely to overcome these drawbacks and provide a better representation of the relationship between Syrian DOD and the PDO.

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Figure 6. Regressions of ERA-Interim 10 meter horizontal winds (green vectors; m s^{-1}) and cubic wind speed (shading; m^3 s^{-3}) onto standardized PDO index (a) from 2003-2015 and (c) from 1979-2015 and onto the standardized (b) Aqua and (d) Terra DOD indices. Areas where the regressions of the wind speed are significant at the 95% confidence level are dotted and vectors significant at the 90% confidence level are plotted in blue (t-test).

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Figure 12. Composites of (a) vertically integrated mass-weighted moisture flux (vectors; kg $m^{-1}s^{-1}$) and its magnitude (shading) from the NCEP1 and (b) CAPE (10^4 J/kg) along with 850 hPa winds (m s^{-1}) from the ERA-Interim for the days with Syrian DOD index (Aqua) greater than 1 standard deviation during MAM from 2003-2015. Moisture flux is integrated from surface to 300 hPa. Shading areas significant at the 95% confidence level are dotted.

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	PDO	Niño3.4	P1	P2	LAI1	LAI2	10 m wind
Aqua DOD	**-0.51**	**-0.24**	-0.05	-0.05	**-0.35**	**-0.41**	-0.00
Terra DOD	**-0.50**	**-0.23**	-0.05	-0.06	**-0.32**	**-0.39**	-0.02

Table 2 Same as table 1 but for MAM average from 2003-2015 (or 2014). Coefficients significant at the 95% confidence level are in bold.

	PDO	Niño3.4	P1	P2	LAI1	LAI2	10 m wind
Aqua DOD	**-0.90**	**-0.60**	-0.14	-0.38	**-0.60**	**-0.64**	0.31
Terra DOD	**-0.90**	**-0.60**	-0.10	-0.34	**-0.57**	**-0.61**	0.30

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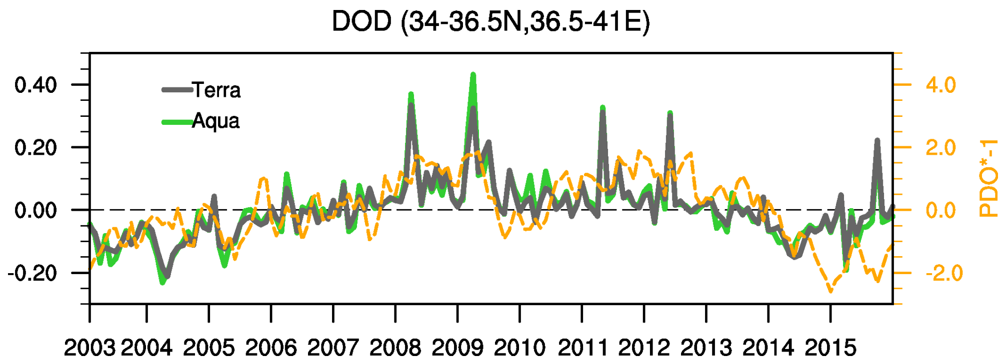

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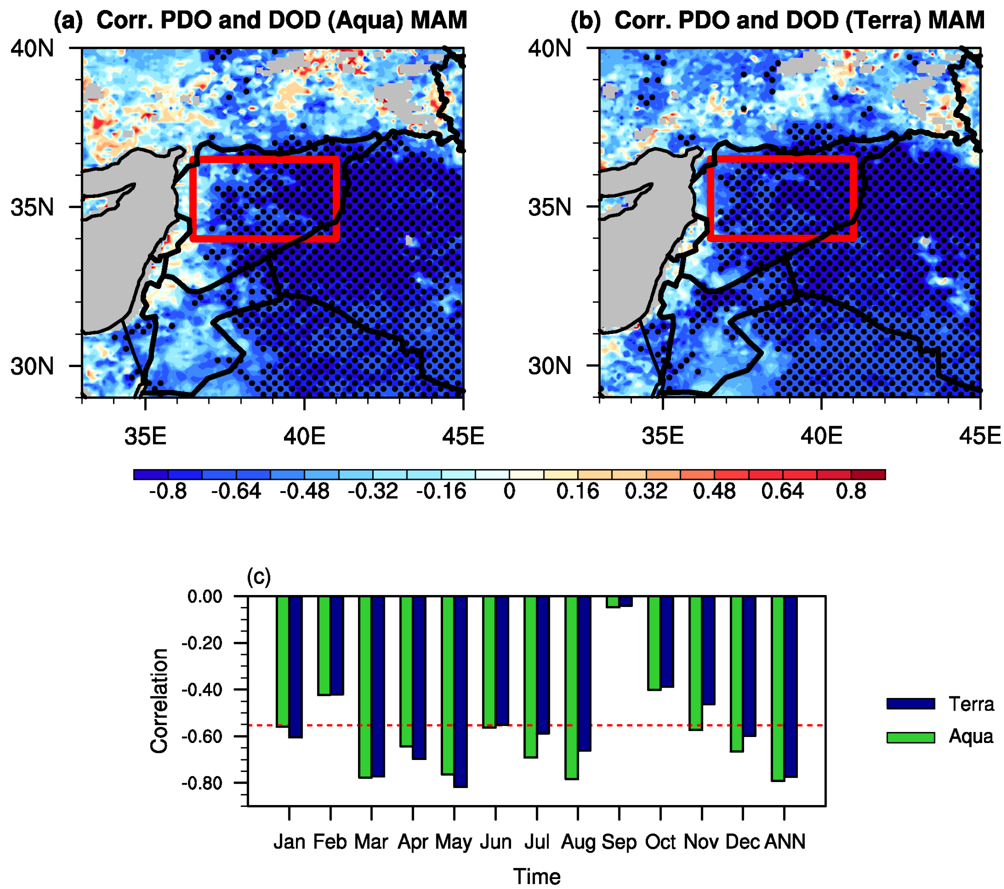

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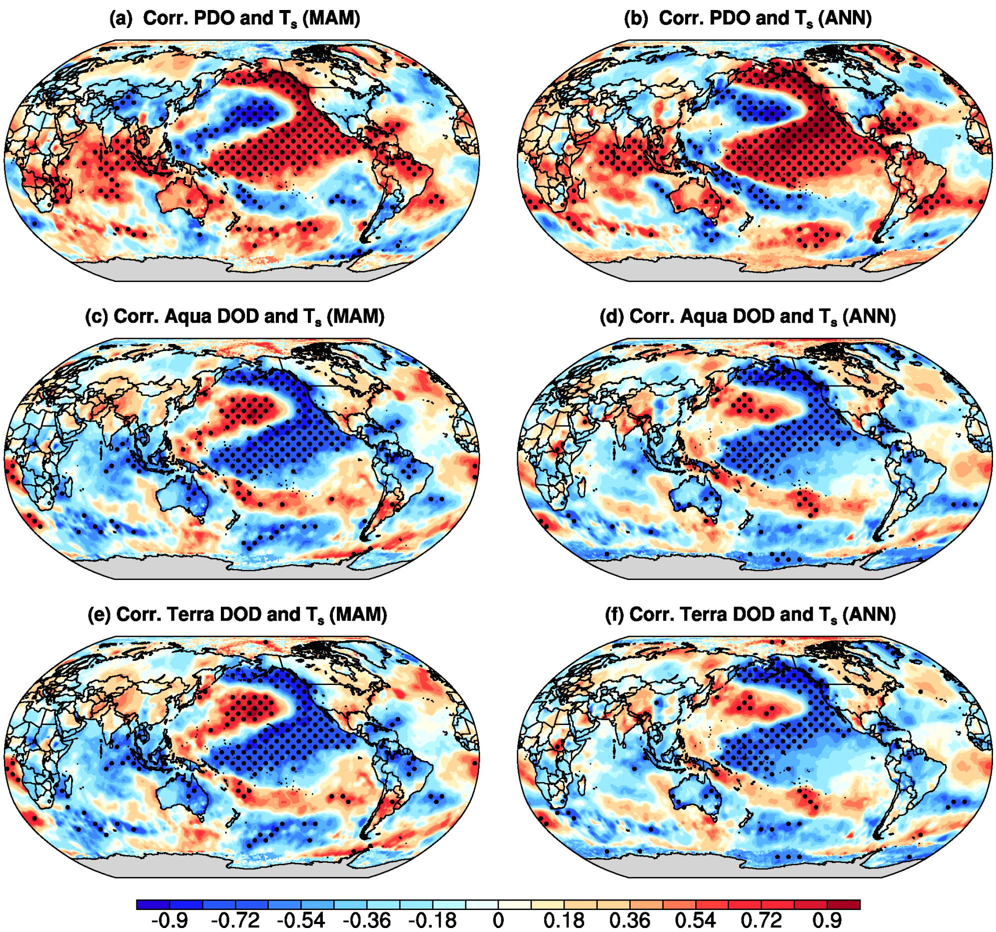

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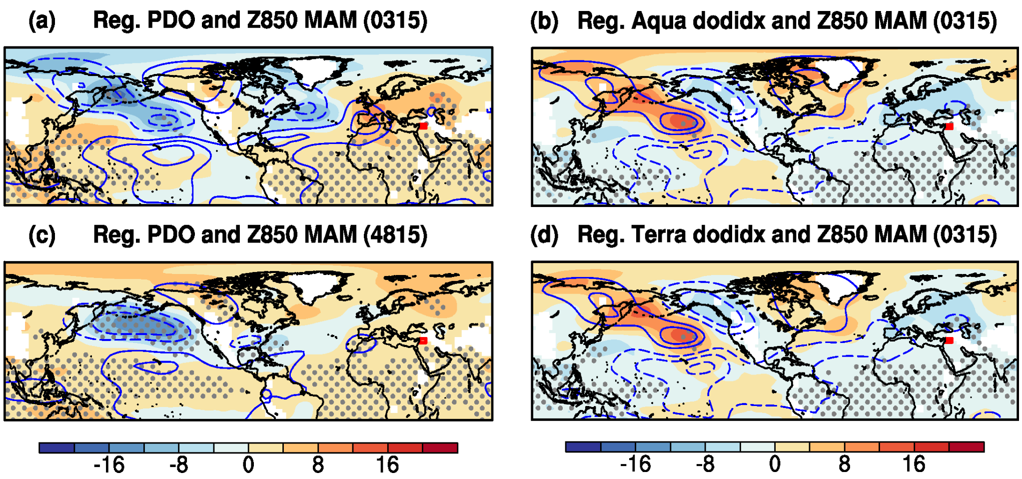

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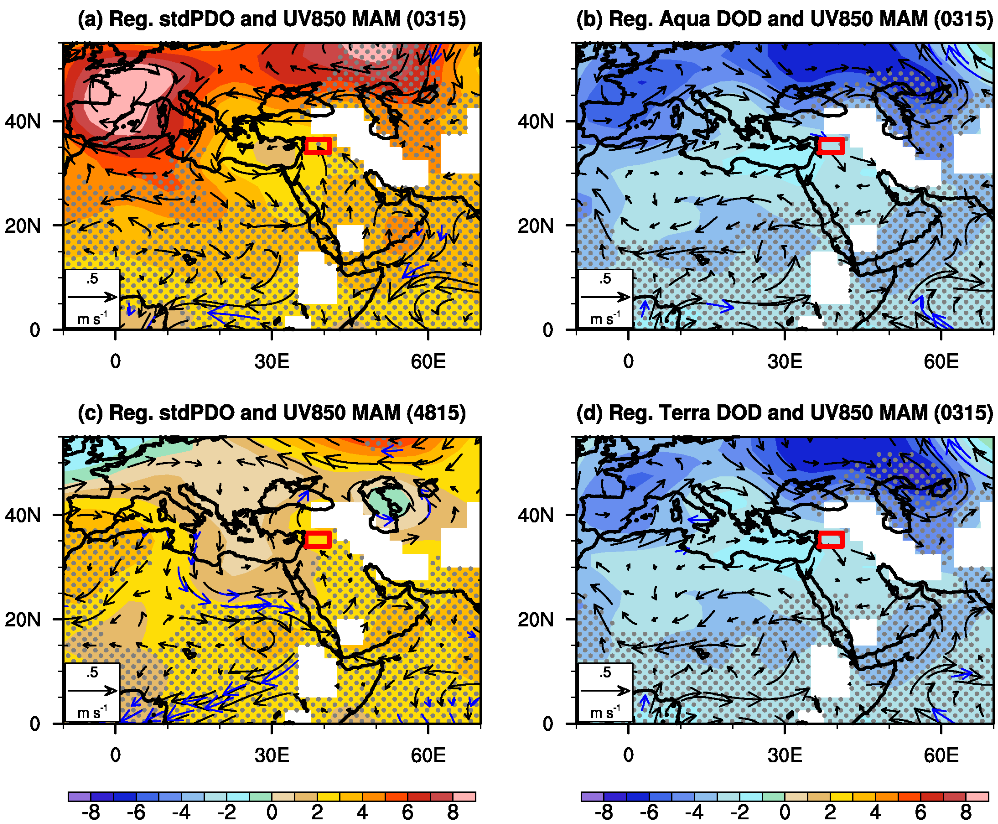

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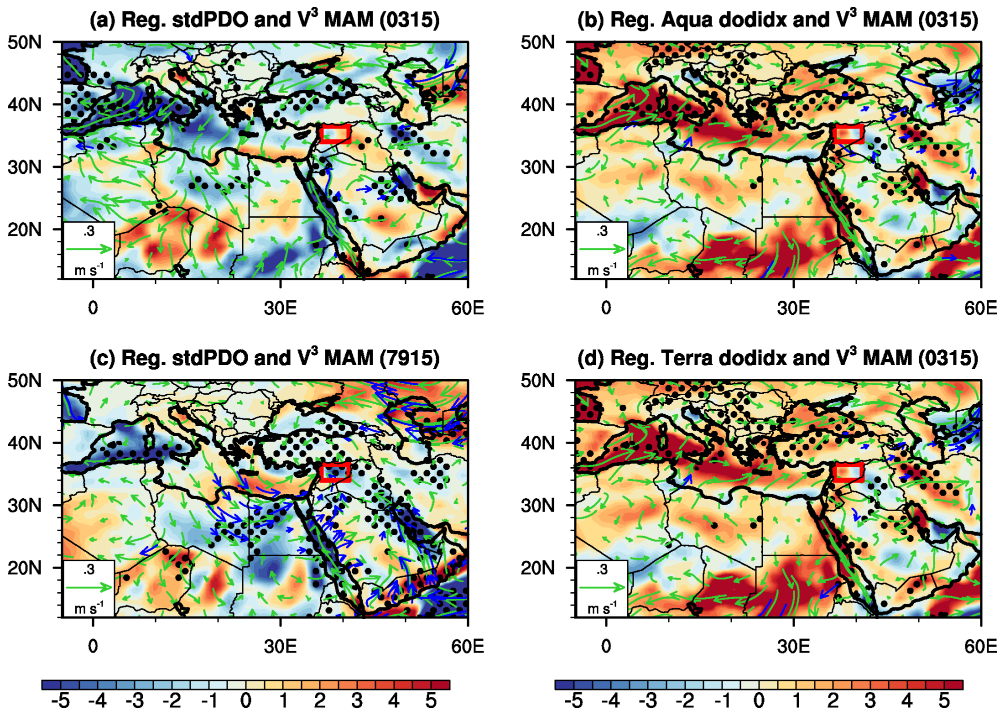

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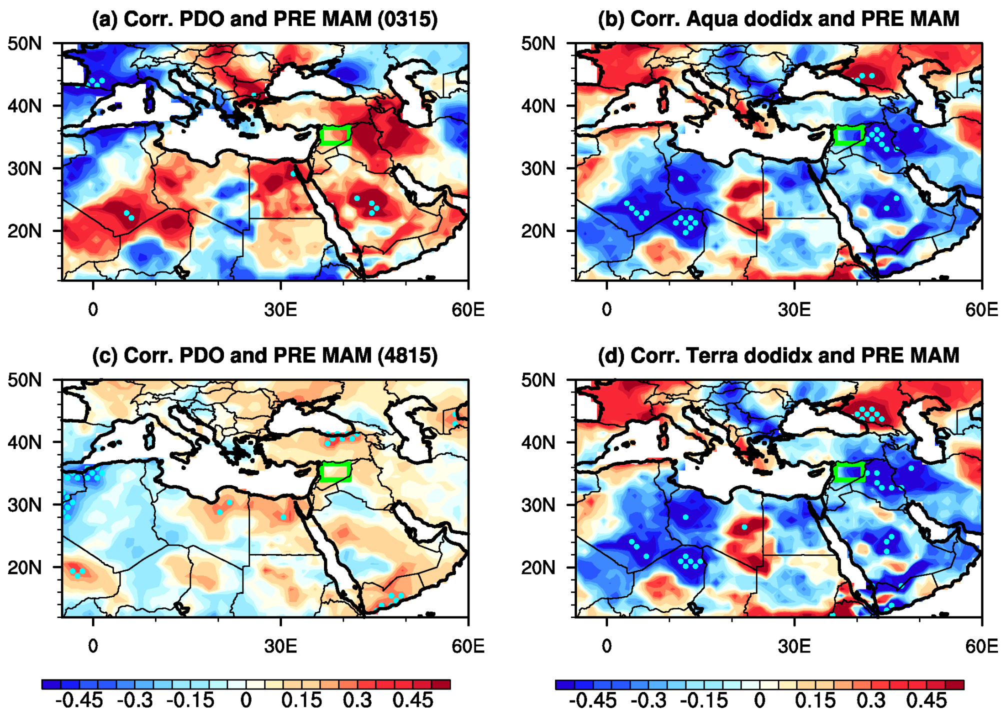

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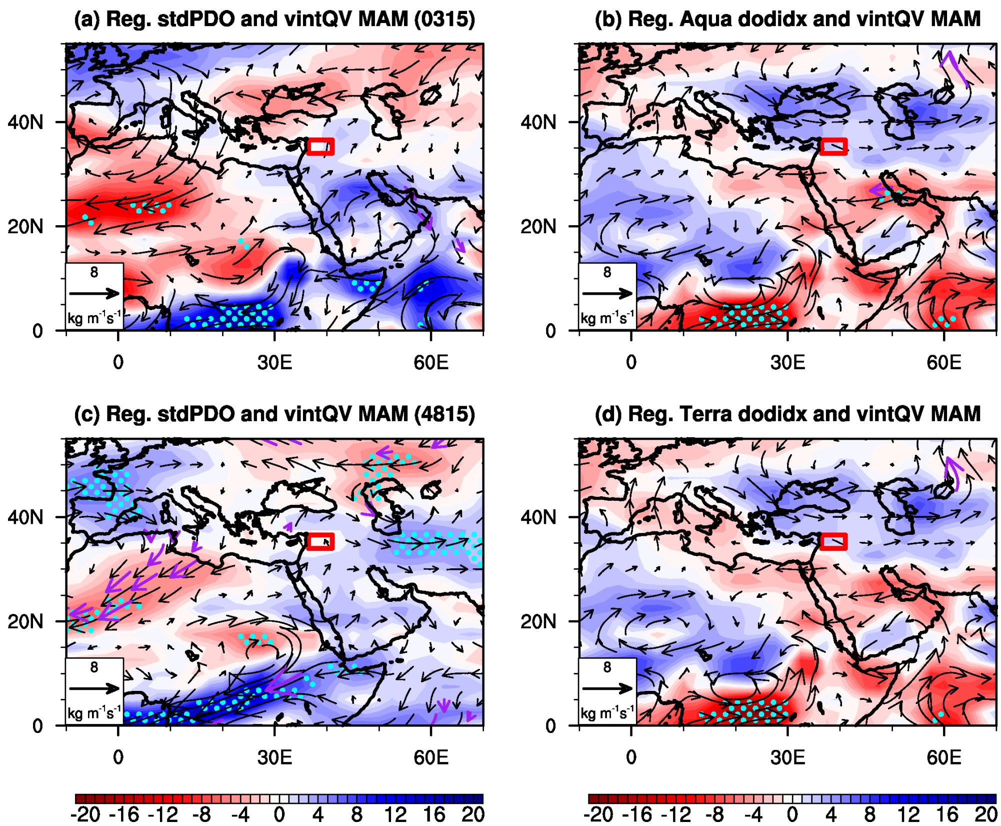

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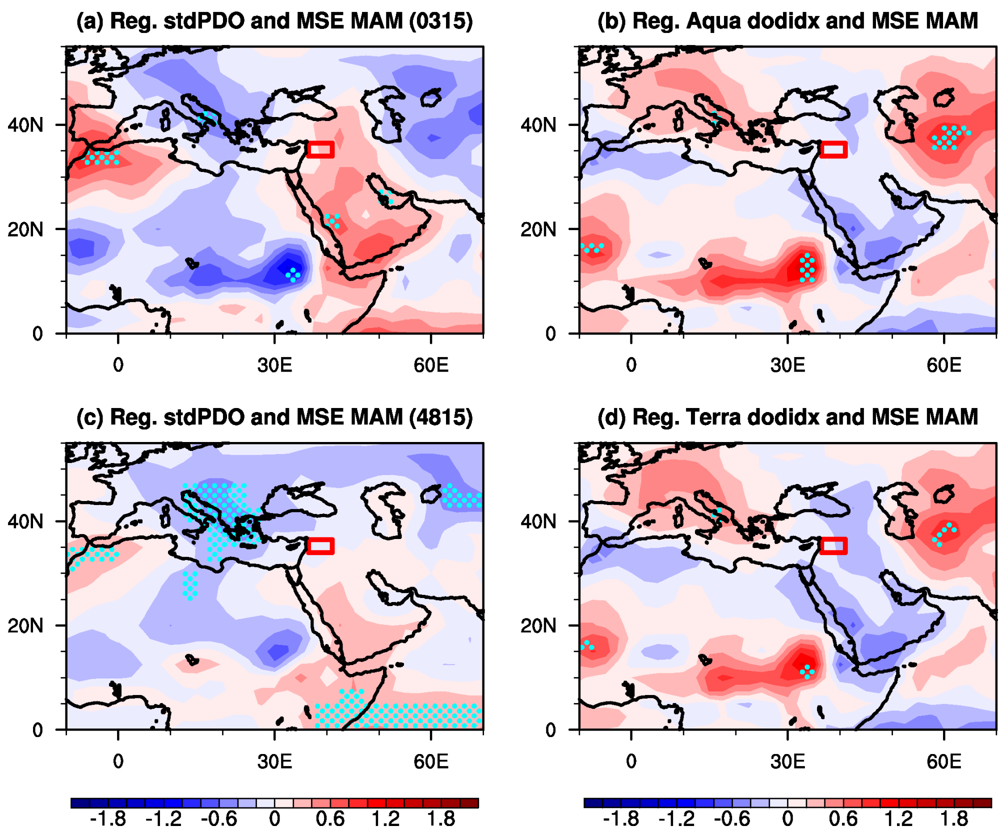

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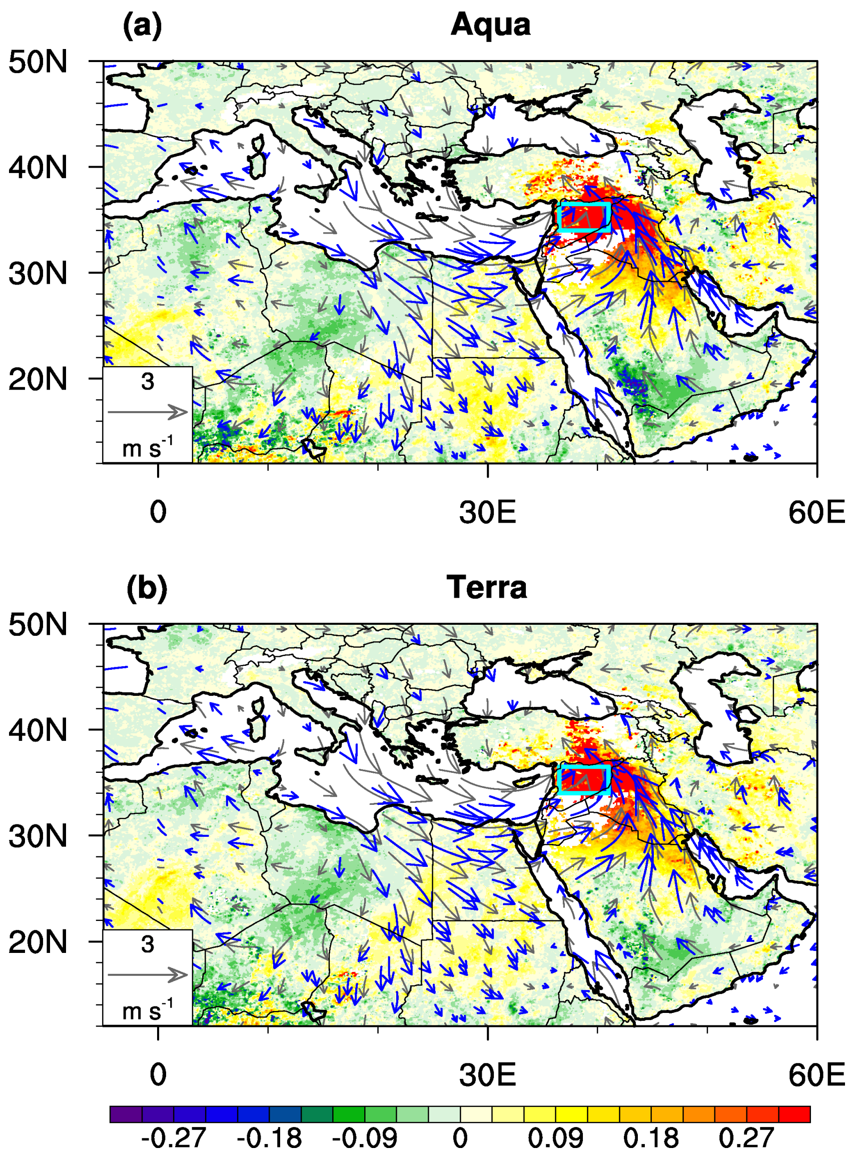

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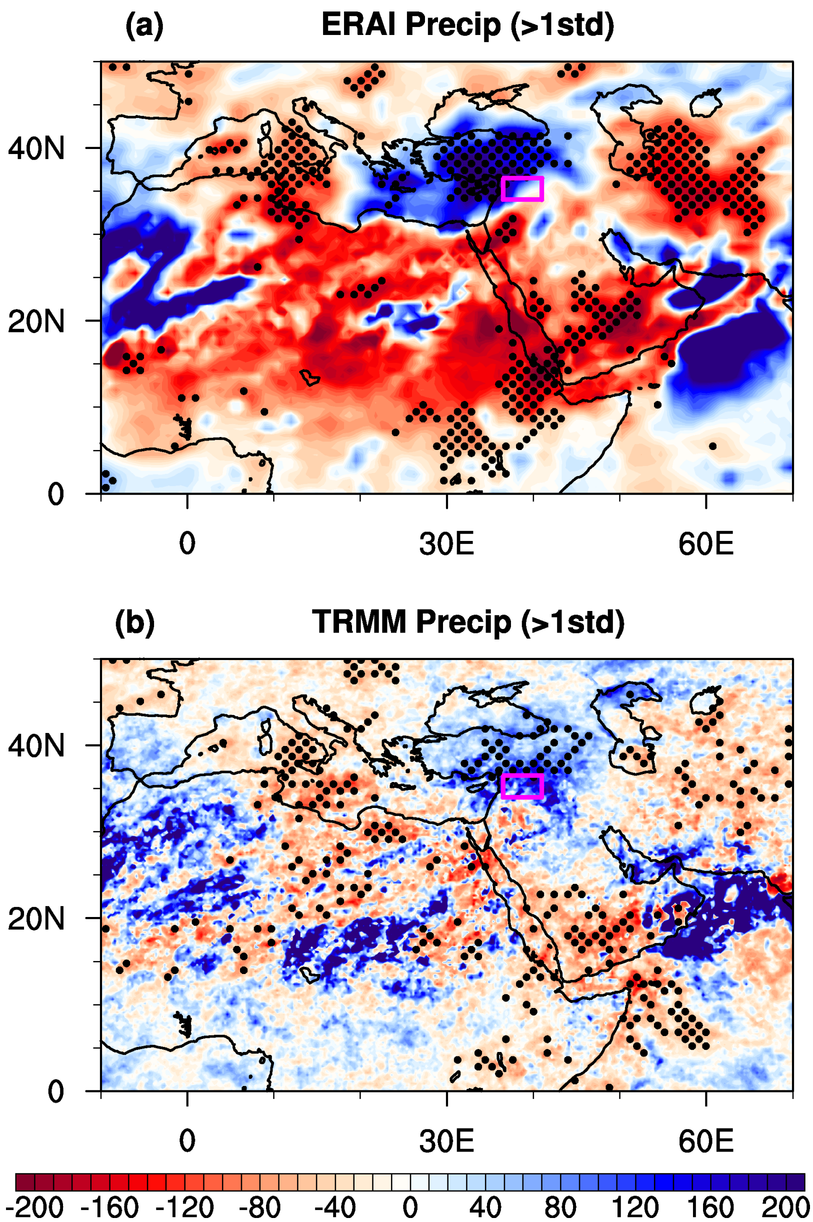

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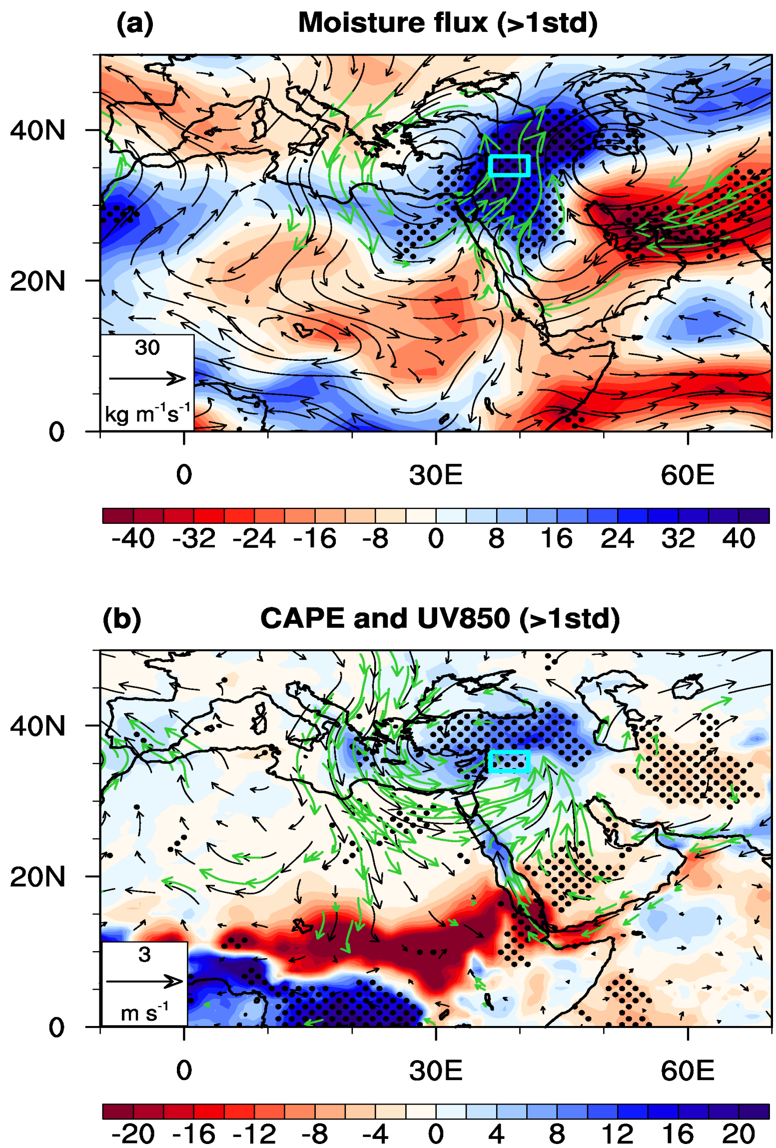

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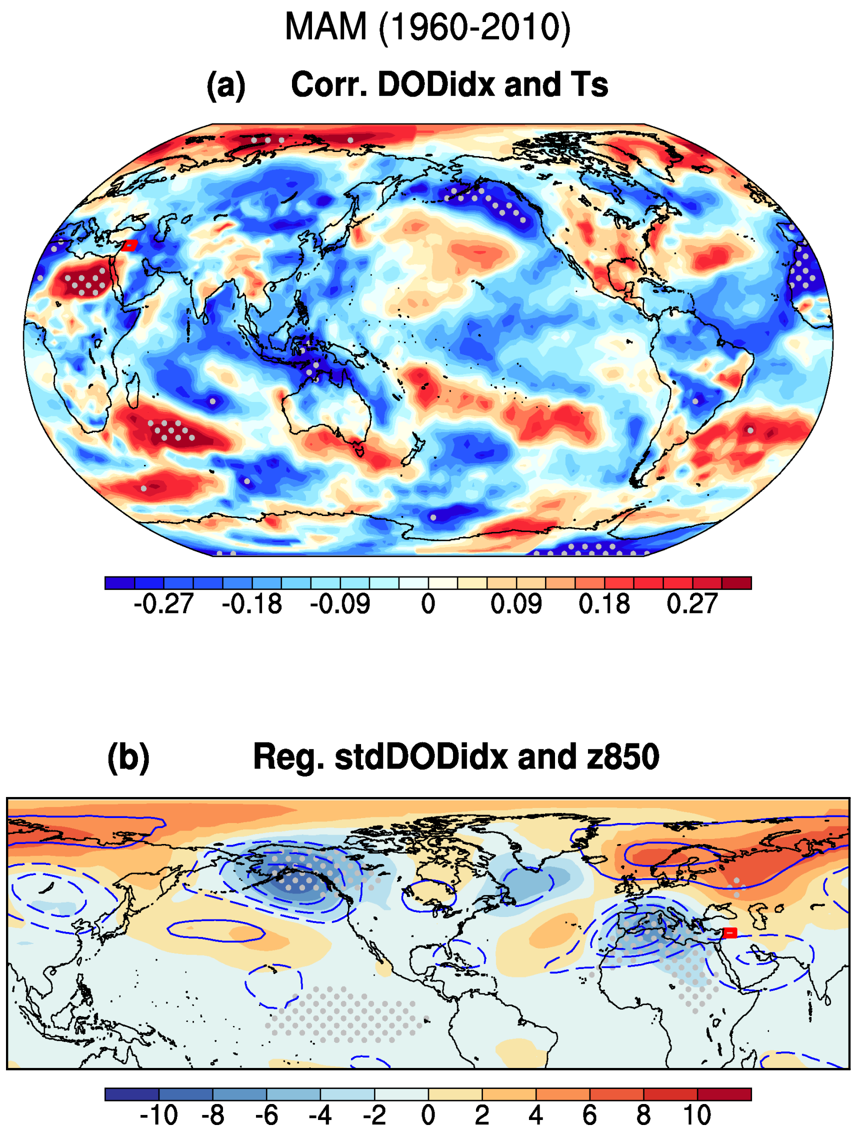

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