

**Response to the comments of Referee #1:**

The paper studied a regional ozone pollution synthetically affected by subtropical high and typhoon system in the Yangtze River Delta region, China. The O<sub>3</sub> pollution has been raised wide attention, and it's an interesting investigation and the manuscript is well written, the discussions are comprehensive, and the results are described clearly. I think that this paper deserves to be published in ACP.

We would like to thank the referee for the valuable and affirmative comments of our manuscript.

**Response to the comments of Referee #2:**

This study discussed the combined effects of two synoptic systems on O<sub>3</sub> pollution in eastern China city clusters. Subtropical high was identified as the main cause of O<sub>3</sub> episode, which was also influenced by a typhoon system. The case study was meaningful and to some extent representative of the typical O<sub>3</sub> episodes in summer of eastern and southern China.

However, many scientific and technical problems must be addressed before this paper is reconsidered to be accepted.

We appreciate the referee for the valuable and constructive reviews of our manuscript. We carefully revise the manuscript based on the following comments.

**Scientific problems:**

1. Both VOCs and NO<sub>x</sub> are key precursors of O<sub>3</sub>, the authors only mentioned and analyzed the patterns of NO<sub>2</sub> and CO. It was not enough to explain the spatial characteristics of O<sub>3</sub> pollution. Even so, the authors did not well establish the possible relationships between O<sub>3</sub> and NO<sub>2</sub> and CO. Then, why to show them in Table 2. Suggest to add more O<sub>3</sub> precursors and deepen the discussions.

Thanks for the constructive comment. We agree that NO<sub>x</sub> and VOC are the most important precursors of O<sub>3</sub>, and the possible relationships between O<sub>3</sub>, VOC and NO<sub>x</sub> should be well established. So following the suggestion, the observed VOC records at an urban site in Shanghai (SAES, 31.17°N, 121.43°E) is added, and Section 3.1 is rewritten.

In the new revised manuscript, the time series of VOCs is shown in the new Fig. 3. The brief description of the site and the measurement method for VOC is added on lines 125-131. Meanwhile, in order to better discuss the O<sub>3</sub>-VOCs-NO<sub>x</sub> relationship and reflect the basic characteristics of this O<sub>3</sub> pollution episode, the temporal variations of NO<sub>2</sub> in Shanghai, Hangzhou and Nanjing are also provided in the new Fig. 3. Based on the new Fig. 3, the temporal variations of O<sub>3</sub> and its precursors (NO<sub>2</sub> and VOC), as well as their internal links, are discussed. Please see lines 293-306 in the revised manuscript.

In the rewritten Section 3.1, we focus on the possible relationships of O<sub>3</sub>, VOC and NO<sub>x</sub>. So, the statistic data for CO in Table 2 and the simple discussion in the original manuscript are deleted. We have tried to find more VOC data to deepen our discussion. Unfortunately, only the records in Shanghai are available in this study, because the data of VOC in the Yangtze River Delta region are very limited and hard to get.

2. Page 17, lines 410-413, the explanation to biases of O<sub>3</sub> and NO<sub>2</sub> were not convincing. The first sentence was meaningless. Although the second sentence described a common sense, it was not responsible for the discrepancies between simulation and observation. Note that the real conditions of O<sub>3</sub> photochemistry are also non-linear.

Thanks for the constructive comment. We carefully analyze the biases of O<sub>3</sub> and NO<sub>2</sub>, and rewrite the words on lines 410-413 in the original manuscript.

In the new revised manuscript, we present three causes resulting in the biases, including the uncertainties in emissions of ozone precursors ("higher estimation of NO<sub>x</sub> emission in Shanghai leads to higher NO<sub>2</sub> and lower O<sub>3</sub> predictions, while lower NO<sub>x</sub> estimations in Nanjing and Hangzhou result in lower NO<sub>2</sub> and higher O<sub>3</sub> modeling results"), the overestimations in WS<sub>10</sub> and the negative biases in T<sub>2</sub>, and some imperfections in the nonlinear chemical reactions coupled in CMAQ (especially the nocturnal chemistry). Please see lines 451-461 of the new revised manuscript.

3. Throughout the paper, the authors did not say why to select these four cities? Particularly for Wuxi, it was not a provincial capital or an extremely polluted city according to the information the authors provided.

Thanks for the constructive comment.

In the original manuscript, the meteorological and air quality observation data at the observation sites in Shanghai (SH, 31.40°N, 121.46°E), Hangzhou (HZ, 30.23°N, 120.16°E), and Nanjing (NJ, 32.00°N, 118.80°E) are used to validate the reliability of simulation by WRF/CMAQ. Unfortunately, we cannot get the meteorological records at the sites in Wuxi (31.62°N, 120.27°E). So, only the observed air quality data are adopted to evaluate the performance of CMAQ. Additionally, to reveal the roles of the individual physical and chemical processes involved in O<sub>3</sub> formation, we present the results of the IPR analysis for the typical cities such as Shanghai, Nanjing and Hangzhou. These cities (Shanghai, Nanjing, Hangzhou, and Wuxi) are all highly urbanized and industrialized, and suffer from severe O<sub>3</sub> pollution. For Shanghai, Nanjing and Hangzhou, they are provincial capitals and typical mega cities in YRD, and they also represent the cities in Southeast Coastal Region (SCR), Northwestern Inland Region (NIR), and Central Inland Region (CIR). For Wuxi, it is located between Shanghai and Nanjing, and also close to the Taihu Lake. So, we chose these cities for model validation and further analysis.

We are sorry that we did not emphasize why to select these cities in the original manuscript. In the new revised manuscript, we provide more information in Methodology to indicate our consideration. We rewrite the section "2.1 Observed meteorological and chemical data", and revise some words in the section "2.3 Integrated Process Rate (IPR) analysis method" and the section "2.4 Evaluation method". The comparison between the observed and the simulated O<sub>3</sub> concentrations in Wuxi is just a supplement for the evaluation of CMAQ. On account that the comparisons in Shanghai, Nanjing and Hangzhou are enough to prove the good performance of our simulations, the comparison for O<sub>3</sub> in Wuxi is deleted. Please see new figure 6 and lines 440-467 in the revised manuscript.

4. Why were the modellings of wind speed and direction not validated with the observation data? They are closely related to the transport processes.

Thanks for the constructive comment. We agree that the wind components "are closely related to the transport processes", and the modeling results of wind should be "validated with the observation data".

In the new revised manuscript, the validation for the modeling results of wind speed and direction is added in Section 4.1. Please see lines 417-425 and Table 3 of the revised manuscript.

5. The northwest movement of high O<sub>3</sub> centers from 10 to 12 Aug. was not obvious. On the basis of Fig. 6, the center located at around 34N, 34N and 33N on these three consecutive days. Clarify and revise the discussions accordingly.

Thanks for the constructive comment. According to the suggestion, in order to clarify the northwest movement of high O<sub>3</sub> centers from 10 to 12 Aug., we replace the old Fig. 6 in the original manuscript by the new Fig. 7. The new Fig. 7 in the revised manuscript presents the daytime vertical wind velocity and the vertical distribution of O<sub>3</sub> from 116.5°E to 122.9°E along the latitude of 31.40°N (where Shanghai is located). Fig. 7d-f clearly show that the high value center of O<sub>3</sub> concentrations (above 90 ppb) moves westwards during August 10-12, implying that the peripheral circulation of Typhoon Utor can drive the air from the coastal areas to the inland areas. Accordingly, we revise the discussions. Please see lines 476-507 of the new revised manuscript.

6. What was meant when vertical diffusion was used? How to explain diffusion process aggravated O<sub>3</sub> pollution? Whether it was accurate to use vertical diffusion if the downward air flow transported high O<sub>3</sub> to the surface.

It is clear that due to the downward airflows (the upper parts of Fig. 7 in the new revised manuscript) dominated by the subtropical high and the typhoon system, the YRD region is under the stable, fine and hot weather condition, and the high concentrations (> 60 ppb) of ozone usually appear from the surface to 1.5 km height. However, as shown in Fig. 7 and 8 (revised in the new manuscript), the near-surface vertical velocity is much lower than that of higher altitudes. Especially in the planetary boundary layer near cities (< 1 km), lots of zero-velocity lines appear near the ground (Fig.8). This phenomenon may be related with the upward airflow caused by Urban Heat Islands. Thus, the maximum centers of O<sub>3</sub> occur near the surface below 500 m, and the vertical diffusion process plays a more important role in the accumulation of surface O<sub>3</sub> than the advection process. The role of vertical diffusion process in the O<sub>3</sub> episode is similar to that reported by Zhu et al. (2015).

Reference: Zhu, B., Kang, H. Q., Zhu, T., Su, J. F., Hou, X. W., and Gao, J. H.: Impact of Shanghai urban land surface forcing on downstream city ozone chemistry, *J Geophys Res-Atmos*, 120, 4340-4351, 10.1002/2014JD022859, 2015.

The above explanations are also added in the new revised manuscript to avoid the confusion from readers. Please see lines 476-558.

7. Page 23, lines 541-549, evidences need to be provided to confirm whether NJ was influenced by typhoon.

Thanks for the constructive comment.

Shanghai and Hangzhou are the cities more close to the sea, so they are easily affected by the typhoon system. Before August 13, O<sub>3</sub> concentrations in these cities exceed the national O<sub>3</sub> standard because of the synthetic effects of subtropical high and typhoon system. After August 13, the YRD region is totally under the control of the typhoon system, and thereby the hot weather is relieved and the O<sub>3</sub> pollution is mitigated. But for Nanjing, because it is far away from the coastal areas, it is hardly affected by the downward flow in the typhoon periphery. As shown in Fig.2 and Fig.8, the concentration of O<sub>3</sub> in Nanjing does not exceed the national O<sub>3</sub> standard. The increase of O<sub>3</sub> in Nanjing on August 12 (Fig.8c) should mainly be caused by the local photochemical reactions because the vertical movement below 2 km above Nanjing is dominated by upward airflows. In Fig. 9c, the IPR analysis also shows that Nanjing is not easily affected by the peripheral circulation of Typhoon Utor. After August 14 (the typhoon system significantly weakens the high pressure system), however, the decrease of O<sub>3</sub> concentration in Nanjing is influenced by the surface wind brought by the typhoon system.

According to this suggestion, we add more evidences in Fig. 8, including the temporal variations of the vertical wind velocity and the vertical distribution of O<sub>3</sub> above Hangzhou and Nanjing during August 7 to 12 2013. We also add some words (as shown above) to deepen our discussion. Please see lines 538-543 and 611-620 of the new revised manuscript.

8. In the section of process analyses, the observed and simulated weather conditions in terms of spatial and temporal patterns should be provided to aid the analyses.

Thanks for the constructive comment. We have added the relevant analysis in Section 4.3. Please see lines 562-610 of the new revised manuscript.

9. Page 25, lines 586-591, why sea-land breeze contributed positively to land O<sub>3</sub> and negatively to oceanic O<sub>3</sub>? More supporting information are needed, such as the distribution map of O<sub>3</sub> and wind fields.

Thanks for the constructive comment. On account of the high-pressure system and so-caused sinking airflows in the YRD region, the background wind is relatively weak in comparison to the local atmospheric circulation, thus the sea breeze can easily bring more generated O<sub>3</sub> to the shore. We add some discussion in Section 4.3.2. Please see lines 662-665 of the new revised manuscript.

10. Subtropical high is a common weather system dominating East Asia region in summer. Why did it only cause O<sub>3</sub> episode in this period? Whether the effect was strengthened by the typhoon? On one hand, in the typhoon periphery, the strong downward flow stimulates O<sub>3</sub> formation and suppresses air pollutants diffusion. On the other hand, had typhoon brought dirty air to the region from inland cities? Overall, the process analyses need to be more comprehensive. During the episode, when the subtropical high dominated and when typhoon dominated? Their combined

effect was promotion or offset?

Thanks for the constructive comment.

Although Western Pacific subtropical high (WPSH) is a normal influential system for the East China region in summer, but the WPSH is stronger and extends much farther west than normal in this period. The anomaly of the WPSH is the major and direct impacting factor for the hot weather (Peng et al., 2014). Besides, this O<sub>3</sub> episode is also influenced by the typhoon system. In the typhoon periphery, the strong downward flow may suppress air pollutants diffusion and cause contamination accumulation. It is not obvious that the typhoon brought dirty air to the region from inland cities. Instead, the clean marine inlet air brought by typhoon is advantageous to the dilution and diffusion of pollutants after the O<sub>3</sub> episode. We add the above explanation in the new revised manuscript.

To make the process analyses more comprehensive and answer the questions of reviewers, we rewrite the section 4.3.2 and new Fig. 11 in the new revised manuscript. Fig. 11 shows the differences of the contributions of main processes between the period of August 7-9 and August 10-12, which can quantitatively evaluate the role of the typhoon system in this severe high O<sub>3</sub> episode. It is clear that the abnormally strong WPSH dominated during the whole O<sub>3</sub> episode. Meantime, from August 10 to 12, the YRD is influenced by the typhoon system as well, because the changes in the contributions of VDIF, CHEM, DDEP, and TADV exhibit a similar spatial pattern with the high values mostly concentrating in the southeast coastal areas. Their combined effect is promoted during August 10 to 12.

Reference: Peng, J. B.: An Investigation of the Formation of the Heat Wave in Southern China in Summer 2013 and the Relevant Abnormal Subtropical High Activities, *Atmospheric & Oceanic Science Letters*, 7, 286-290, 2014.

**Technical problems:**

1. Too many grammatical errors in this paper, e.g. 1, line 15 “is detected”, line 16, exceeding”, “reaches”, line 22 “abnormal strong”, line 25 “worse air pollution”. I cannot list all of them. Strongly suggest to correct the errors with the aid of a professional language correcting company. Sorry for these grammatical errors in the original manuscript. The errors listed above are corrected as follows.

The words “is detected” on line 14 of the original manuscript are revised to “was detected”. Please see line 13 in the new revised manuscript.

The word “exceeding” on line 16 of the original manuscript is revised to “exceeded”. Please see line 15 in the new revised manuscript.

The words “abnormal strong” on line 22 of the original manuscript are revised to “abnormally strong”. Please see line 21 in the new revised manuscript.

The words “worse air pollution” on line 25 of the original manuscript are revised to “worse air

quality". Please see line 24 of the new revised manuscript.

Additionally, the authors asked an English native speaker (named Josh Powell from Georgetown University) to help improving the English.

2. Define the abbreviations at their first appearances, e.g. WRF-CMAQ.

Thanks for the constructive comment.

As suggested above, "CMAQ" on line 27 (where "CMAQ" occurs for the first time) of the original manuscript is replaced by "the Community Multi-scale Air Quality (CMAQ) Model". Please see lines 26-27 in the new revised manuscript.

The words "With the aid of the WRF/CMAQ" on lines 96 (where "WRF/CMAQ" appears for the first time) in the original manuscript are rewritten as "The WRF/CMAQ model system, which consists of the Weather Research and Forecasting (WRF) model and the Community Multi-scale Air Quality (CMAQ) Model, were used to reveal the exact formation mechanism." in the new revised manuscript. Please see lines 98-101 in the new revised manuscript.

"(LST)" on line 146 (where "LST" appears for the first time) in the original manuscript is changed to "(local standard time, LST)". Please see line 164 in the new revised manuscript.

"PBL" on line 154 (where "PBL" appears for the first time) in the original manuscript is revised to "planetary boundary layer". Please see lines 172-173 in the new revised manuscript.

"USGS" on line 157 (where "USGS" appears for the first time) in the original manuscript is replaced by "United States Geological Survey (USGS)". Please see line 176 in the new revised manuscript.

The sentence "The initial meteorological fields and boundary conditions are from NCEP FNL global reanalysis data with  $1^\circ \times 1^\circ$  resolution" on lines 160-161 (where "NCEP" appears for the first time) in the original manuscript is rewritten as "The initial meteorological fields and boundary conditions are from  $1^\circ$  resolution global reanalysis data provided by National Center for Environmental Prediction (NCEP)". Please see lines 178-180 in the new revised manuscript.

"(UTC)" on line 227 (where "UTC" appears for the first time) in the original manuscript is changed to "(Universal Time Coordinated, UTC)". Please see line 238 in the new revised manuscript.

"WRF-CMAQ" on line 642 of the original manuscript is revised to "WRF/CMAQ". Please see line 735 of the new revised manuscript. The definition has been given at its first appearances (on lines 98-101 in the new revised manuscript).

3. Page 4, line 94, what is the knowledge gap?

“Knowledge gap” means the principle that we do not know and need to be studied. The phrase “to fill the knowledge gap” was also used in some other papers (Ding et al., 2013; Xie et al., 2016). Here, the authors wanted to express that it is worth to investigate how the subtropical high and the typhoon system affect the formation of this regional O<sub>3</sub> pollution, and the results will help us to understand the important factors impacting O<sub>3</sub> formation from the regional scale.

To avoid unnecessary misunderstanding, the words “To fill the knowledge gap and better understand the important factors impacting O<sub>3</sub> formation from the regional scale, we perform an observational analysis to identify the temporal and spatial characteristics of the episode. With the aid of the WRF/CMAQ as well as the Integrated Process Rate analysis (IPR) coupled within CMAQ, numerical simulations are conducted to provide qualitative and quantitative analysis on the contributions of individual atmospheric processes.” on lines 94-98 of the original manuscript are revised to “To better understand the important factors impacting O<sub>3</sub> formation from the regional scale, we investigated the exact roles of these two typical weather systems in this pollution episode by using observational analysis and numerical simulations. The observational analysis was performed to identify the temporal and spatial characteristics of the episode. The WRF/CMAQ model system, which consists of the Weather Research and Forecasting (WRF) model and the Community Multi-scale Air Quality (CMAQ) Model, were used to reveal the exact formation mechanism. With the aid the Integrated Process Rate (IPR) analysis coupled in CMAQ, the qualitative and the quantitative analysis on the contributions of individual atmospheric processes were conducted as well.”. Please see line 95-103 of the new revised manuscript.

4. Past tense was suggested for the Introduction section, except for the common senses.

Thanks for the constructive comment. As suggested above, several sentences in the Introduction section are revised to use the past tense.

The words “there is” on line 92 of the original manuscript are revised to “there was”. Please see line 93 in the new revised manuscript.

The words “may be” on line 93 of the original manuscript are revised to “might be”. Please see line 94 in the new revised manuscript.

The words “we perform” on line 95 and “are conducted” on line 97 of the original manuscript are deleted. “..., we investigated ...”, “The observational analysis was performed to ...”, “..., were used ...” and “... were conducted” are added in the new revised manuscript. Please see lines 97-103 of the revised manuscript.

5. Website references needed to be added for the citations of meteorological and air pollutants data.

Thanks for the constructive comment. The website references for the citations of the meteorological and air pollutants data are added in the new revised manuscript. For the meteorological data, “The weather charts for East Asia are accessible from Korea Meteorological Administration (<http://www.kma.go.kr/chn/weather/images/analysischart.jsp>)”, and “The hourly

meteorological data at the observation sites of SH (31.40°N, 121.46°E) located in Shanghai, HZ (30.23°N, 120.16°E) in Hangzhou, and NJ (32.00°N, 118.80°E) in Nanjing can be obtained from the University of Wyoming (<http://weather.uwyo.edu/wyoming/>). For the air pollutant data, “The in-situ monitoring data for the hourly concentrations of O<sub>3</sub>, CO, NO<sub>2</sub>, SO<sub>2</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> can be acquired from National Environmental Monitoring Center (NEMC) (<http://106.37.208.233:20035>)”. Please see lines 116-118 and 132-138 in the new revised manuscript.

6. Page 9, lines 236-238, the function of water vapor and the evidence for its more abundance in Shanghai.

Thanks for this constructive comment. In the new revised manuscript, we pay more attention to the chemical relations between O<sub>3</sub> and its precursors. Section 3.1 is rewritten. Water vapor is not discussed, and the words on lines 236-238 of the original manuscript are changed to "It seems that O<sub>3</sub> concentrations are higher in the cities around Shanghai, where the concentrations of O<sub>3</sub> precursors (shown in Table 2) are more adequate as well.". Please see lines 285-287 of the new revised manuscript.

7. Page 17, line 395, WRF/CHEM or WRF-CMAQ?

Sorry for this clerical error. The word “WRF/CHEM” on line 395 of the original manuscript is revised to “WRF/CMAQ”. Please see line 436 of the new revised manuscript.

8. Fig. 6, the legends need to be provided. Positive wind speeds with solid lines mean downward air flow?

Thanks for the constructive comment. The dotted lines show the negative wind speeds and represent downward airflow, while the solid lines show the positive vertical wind speeds and zero vertical velocity. In the new revised manuscript, the legends and the necessary explanation are added to Fig. 7 and 8. Please see lines 512-514 and 548-550 of the new revised manuscript.

9. Conclusion needs to be reorganized after the revision of the whole paper.

As suggested above, the conclusion is reorganized and revised after the revision of the whole paper. Please see lines 705-740 of the new revised manuscript.



1 **Integrated studies of a regional ozone pollution synthetically**  
2 **affected by subtropical high and typhoon system in the**  
3 **Yangtze River Delta region, China**

4 Lei Shu<sup>†</sup>, Min Xie<sup>†\*</sup>, Tijian Wang<sup>†2\*</sup>, ~~Da Gao~~, Pulong Chen<sup>†</sup>, Yong Han<sup>†</sup>, Shu Li<sup>†</sup>, Bingliang  
5 Zhuang<sup>†</sup>, Mengmeng Li<sup>†</sup>, ~~Da Gao~~<sup>†</sup>

6 <sup>†</sup>School of Atmospheric Sciences, Nanjing University, Nanjing, China

7 <sup>2</sup>CMA-NJU Joint Laboratory for Climate Prediction Studies, Institute for Climate and Global  
8 Change Research, School of Atmospheric Sciences, Nanjing University, Nanjing, China

9 [School of Atmospheric Sciences, CMA-NJU Joint Laboratory for Climate Prediction Studies,](#)  
10 [Jiangsu Collaborative Innovation Center for Climate Change, Nanjing University, Nanjing 210023,](#)  
11 [China](#)

12 -----  
13 \*Corresponding author, +86-25-89685302; [E-mail address: \(minxie@nju.edu.cn\)](mailto:minxie@nju.edu.cn); and [Tijian Wang \(tjwang@nju.edu.cn\)](mailto:tjwang@nju.edu.cn)

16 **Abstract:** Severe high ozone (O<sub>3</sub>) episodes usually have close relations to synoptic systems. A  
17 regional continuous O<sub>3</sub> pollution episode ~~is-was~~ detected over the Yangtze River Delta (YRD)  
18 region in China during August 7-12, 2013, in which the O<sub>3</sub> concentrations in more than half of the  
19 cities ~~exceeding-exceeded~~ the national air quality standard. The maximum hourly concentration of  
20 O<sub>3</sub> ~~reaches-reached~~ 167.1 ppb. By means of the observational analysis and the ~~WRF/CMAQ~~  
21 numerical simulation, the characteristics and the essential impact factors of the typical regional O<sub>3</sub>  
22 pollution ~~is-are~~ integratedly investigated. The observational analysis shows that the atmospheric  
23 subsidence dominated by Western Pacific subtropical high plays a crucial role in the formation of  
24 high-level O<sub>3</sub>. The favorable weather conditions, such as extremely high temperature, low relative  
25 humidity and weak wind speed, caused by the abnormally ~~ly~~ strong subtropical high are responsible  
26 for the trapping and the chemical production of O<sub>3</sub> in the boundary layer. In addition, when the  
27 YRD cities at the front of Typhoon Utor, the periphery circulation of typhoon system can enhance  
28 the downward airflows and cause ~~worse-more-seriousworse~~ air ~~pollutionquality~~. But when the  
29 typhoon system weakens the subtropical high, the prevailing southeasterly surface wind leads to

带格式的: 无孤行控制

带格式的: 无孤行控制

30 the mitigation of the O<sub>3</sub> pollution. The Integrated Process Rate (IPR) analysis incorporated in [the](#)  
31 [Community Multi-scale Air Quality \(CMAQ\) Model](#) is applied to further illustrate the combined  
32 influence of subtropical high and typhoon system in this O<sub>3</sub> episode. The results show that the  
33 vertical diffusion (VDIF) and the gas-phase chemistry (CHEM) are two major contributors to O<sub>3</sub>  
34 formation. During the episode, the contributions of VDIF and CHEM to O<sub>3</sub> maintain the high  
35 values over ~~10 ppb/h in Shanghai, Hangzhou, and Nanjing~~ [the YRD region](#). On August 10-11-12,  
36 the cities close to the sea are apparently affected by the typhoon system, with the contribution of  
37 VDIF increasing to 28.45 ppb/h in Shanghai and 19.76 ppb/h in Hangzhou. [In contrast, the cities](#)  
38 [in the northwest inland area far away from the sea are generally under the control of the](#)  
39 [subtropical high and can hardly be affected by the periphery circulation of typhoon system](#). When  
40 ~~the YRD region is under the control of~~ the typhoon system [significantly weakens the subtropical](#)  
41 [high](#), the contribution values of all individual processes decrease to a low level in all [YRD](#) cities.  
42 These results provide an insight for the O<sub>3</sub> pollution synthetically impacted by the Western Pacific  
43 subtropical high and the tropical cyclone system.

44 **Keyword:** Ozone; subtropical high; typhoon; the Yangtze River Delta region; heat wave

45

## 46 1. Introduction

47 Ground-level ozone (O<sub>3</sub>) is a secondary air pollutant generated by a series of complicated  
48 photochemical reactions involving nitrogen oxides (NO<sub>x</sub>) and hydrocarbons (HC) (Crutzen, 1973;  
49 Sillman, 1999; Jenkin et al., 2000; Wang et al., 2006b; Xie et al., 2014; 2016b). Severe O<sub>3</sub>  
50 pollution events usually occur in the presence of sunlight and under favorable meteorological  
51 conditions, with the abundance of O<sub>3</sub> precursors (NO<sub>x</sub> and HC) (Wang et al., 2006b). These O<sub>3</sub>  
52 pollutions in troposphere can deteriorate the air quality, and thereby cause adverse effects on  
53 human health and vegetation (Feng et al., 2003; Fann and Risley, 2013; Landry et al., 2013).  
54 Consequently, the formation mechanism and the integrated prevention of O<sub>3</sub> pollution are of great  
55 concern in many megacities all over the world (Xie et al., 2016b).

56 Over the past decades, along with the rapid industrial and economic development, many areas  
57 in China have been suffering from high levels of O<sub>3</sub> pollution. Especially in the most economically  
58 vibrant and densely populated areas, such as the Yangtze River Delta (YRD) region, the Pearl  
59 River Delta (PRD) region, and the Beijing-Tianjin-Hebei (BTH) area, the severe O<sub>3</sub> pollution

60 episode has frequently occurred (Lam et al., 2005; Wang et al., 2006b; An et al., 2007; Chan and  
61 Yao, 2008; Duan et al., 2008; Jiang et al., 2008; Zhang et al., 2008; Guo et al., 2009; Shao et al.,  
62 2009; Ma et al., 2012) , and the background air pollutant concentrations have steadily increased  
63 (Chan and Yao, 2008; Zhang et al., 2008; Tang et al., 2009; Wang et al., 2009a; Ma et al., 2012;  
64 Liu et al., 2013). Many studies on the O<sub>3</sub> pollution, including satellite data analyses, field  
65 experiments, and model simulations, have been carried out over China in order to investigate the  
66 temporal and spatial characteristics of surface photochemical pollutions (Lu and Wang, 2006;  
67 Wang et al., 2006a; Tu et al., 2007; Zhang et al., 2007; 2008; Geng et al., 2008; Tang et al., 2008;  
68 2009; Chen et al., 2009; Han et al., 2011; Ding et al., 2013; Xie et al., 2016b), nonlinear  
69 photochemistry of O<sub>3</sub> and its precursors (Lam et al., 2005; Ran et al., 2009; Liu et al., 2010; Li et  
70 al., 2011; Xie et al., 2014), interactions between O<sub>3</sub> and aerosols (Lou et al., 2014; Shi et al., 2015),  
71 the effects of urbanization on O<sub>3</sub> formation (Wang et al., 2007; 2009b; Liao et al., 2015; Li et al.,  
72 2016; Xie et al., 2016a; Zhu et al., 2016), and other essential impact factors (Jiang et al., 2012; Li  
73 et al., 2012; Wei et al., 2012; Liu et al., 2013; Gao et al., 2016).

74 The Yangtze River Delta (YRD) region is a highly developed area of urbanization and  
75 industrialization. With the accelerated economic development and remarkable increase in energy  
76 consumption, the photochemical smog with high level of O<sub>3</sub> concentration is becoming more and  
77 more prominent and frequent, tending to present ~~eonspueuous the regional~~ characteristics of  
78 [regional pollution](#) (Chan and Yao, 2008; Ma et al., 2012; Li et al., 2012). Being located on the  
79 southeastern coast of China, YRD features a typical subtropical monsoon climate and is strongly  
80 affected by the Western Pacific subtropical high in summer. So, high O<sub>3</sub> concentrations are  
81 usually observed in late spring and summer by in-situ monitoring (Ding et al., 2013; Xie et al.,  
82 2016b). Severe high O<sub>3</sub> episodes usually have close relations to synoptic systems (Huang et al.,  
83 2005; 2006; Wang et al, 2006b; Jiang et al., 2008; Cheng et al., 2014; Hung and Lo, 2015).  
84 Horizontal and vertical transport processes from upwind O<sub>3</sub>-rich air masses as well as poor  
85 atmospheric diffusion conditions can lead to the accumulation of surface O<sub>3</sub> concentrations and  
86 aggravating the photochemical pollution (Wang et al., 2006b). In previous studies on high O<sub>3</sub>  
87 pollution in the YRD region, some researchers have discussed this issue. For example, Jiang et al.  
88 (2012) investigated the spring O<sub>3</sub> formation over East China, and suggested that O<sub>3</sub> concentrations  
89 over the YRD region were transported and diffused from surrounding areas. Li et al. (2014<sup>2</sup>)

90 presented quantitative analysis on atmospheric processes affecting O<sub>3</sub> concentrations in the typical  
91 YRD cities during a summertime regional high O<sub>3</sub> episode, and found that the maximum  
92 concentration of photochemical pollutants was usually related with the process of transportation.  
93 Gao et al. (2016) evaluated the O<sub>3</sub> concentration during a frequent shifting wind period, and  
94 revealed that vertical mixing played an important positive role in the formation of surface O<sub>3</sub>.  
95 However, these investigations only focused on the O<sub>3</sub> formation mechanism for one megacity  
96 (such as Shanghai, Nanjing and Hangzhou, etc.) or just a single station. Up to now, studies on the  
97 process analysis of high ozone episodes over the YRD are quite limited (Li et al., 2012). So, more  
98 studies should pay attention to the typical weather systems and the exact formation mechanism of  
99 the regional O<sub>3</sub> pollution in this region.

100 During August 7-12 2013, there ~~is~~ was a typical regional O<sub>3</sub> pollution episode in the YRD  
101 region, which ~~may~~ might be ~~combinedly~~ synthetically influenced by the Western Pacific  
102 subtropical high and Typhoon Utor. To ~~fill the knowledge gap and~~ better understand the important  
103 factors impacting O<sub>3</sub> formation from the regional scale, ~~we~~ thoroughly investigated the exact roles  
104 of these two typical weather systems in this pollution episode; by using observational analysis and  
105 numerical simulations. ~~we perform an~~ The observational analysis was performed to identify the  
106 temporal and spatial characteristics of the episode. ~~With the aid of~~ The WRF/CMAQ model  
107 system, which consists of the Weather Research and Forecasting-(WRF) model- (WRF) and the  
108 Community Multi-scale Air Quality -(CMAQ) Model (CMAQ), were used to reveal the exact  
109 formation mechanism. ~~With the aid as well as~~ the Integrated Process Rate (IPR) analysis (IPR)  
110 coupled with CMAQ, ~~numerical simulations are were conducted to provide the~~ qualitative and  
111 the quantitative analysis on the contributions of individual atmospheric processes were conducted  
112 as well. ~~The results may be a great help for the prediction and the prevention of high O<sub>3</sub> pollution~~  
113 ~~events.~~ In this paper, the brief description of observational data and model configurations are  
114 shown in Section 2. The detailed observational analysis of air quality and meteorological  
115 conditions are given in Section 3. The evaluation of model performance and the formation  
116 mechanism of O<sub>3</sub> explored by IPR technique are presented in Section 4. In the end, a summary of  
117 main findings is given in Section 5.

118

## 119 2. Methodology

## 2.1 Observed meteorological and chemical data

The weather charts and the observed surface meteorological records are used to analyze the synoptic systems during the episode in August 2013, as well as to evaluate the modeling results of meteorological factors. The weather charts for East Asia are accessible from Korea Meteorological Administration (<http://www.kma.go.kr/chn/weather/images/analysischart.jsp>). The hourly meteorological data at the observation sites of SH (31.40°N, 121.46°E) located in Shanghai, HZ (30.23°N, 120.16°E) in Hangzhou, and NJ (32.00°N, 118.80°E) in Nanjing can be obtained from the University of Wyoming (<http://weather.uwyo.edu/wyoming/>), where 2 m air temperature, 2 m relative humidity, 10 m wind speed and 10 m wind direction are available.

The air quality observational data are used to identify the regional characteristics of the O<sub>3</sub> episode in August 2013 and to validate the model performance for air pollutants. Fifteen cities are selected as the representative research objects to better reflect the status of O<sub>3</sub> pollution over the YRD region. The locations of these cities are shown in Fig. 1b, which contains Shanghai, 8 cities in Jiangsu province (Changzhou, Nanjing, Nantong, Suzhou, Taizhou, Wuxi, Yangzhou, and Zhenjiang), and 6 cities in Zhejiang province (Hangzhou, Huzhou, Jiaxing, Ningbo, Shaoxing, and Zhoushan). The in-situ monitoring data for the hourly concentrations of O<sub>3</sub>, CO, NO<sub>2</sub>, SO<sub>2</sub>, PM<sub>2.5</sub> and PM<sub>10</sub> can be acquired from National Environmental Monitoring Center (NEMC) (<http://106.37.208.233:20035>). The assurance/quality control (QA/QC) procedures for monitoring strictly follow the national standards (State Environmental Protection Administration of China, 2006). The hourly pollutant concentration for a city is calculated as the average of the pollutant concentrations from several national monitoring sites in that city, which can better characterize the pollution level of the city. In order to identify invalid or lacking data, a checking procedure for these data is performed following the work of Chiqueto and Silva (2010). Finally, only less than 0.2% of the primary data are ignored in the calculation. Moreover, the observed data of total VOCs (TVOC) during August 4-10 at an urban site in Shanghai (SAES, 31.17°N, 121.43°E) is also used. They are provided by Shanghai Academy of Environmental Sciences. The sampling height is about 15 m, and individual VOC species are continuously measured every 30 min by two on-line high performance gas chromatograph with flame ionization detector (GC-FID) systems (Chromato-sud airmoVOC C2-C6 #5250308 and airmoVOC C6-C12 #2260308, France). The details for measurement and QA/QC can refer to Wang et al. (2013).

150 [The weather charts and the observed surface meteorological records are used to analyze the](#)  
151 [synoptic systems during the episode. The weather charts for East Asia are accessible from Korea](#)  
152 [Meteorological Administration \(<http://www.kma.go.kr/chn/weather/images/analysischart.jsp>\). The](#)  
153 [hourly meteorological data at the observation sites of SH \(31.40°N, 121.46°E\) located in Shanghai,](#)  
154 [HZ \(30.23°N, 120.16°E\) in Hangzhou, and NJ \(32.00°N, 118.80°E\) in Nanjing can be obtained](#)  
155 [from the University of Wyoming \(<http://weather.uwyo.edu/wyoming/>\), where 2-m air temperature,](#)  
156 [2-m relative humidity, 10-m wind speed and 10-m wind direction are available.](#)

157 [Meteorological and air quality observation data are also used to validate the reliability of](#)  
158 [simulations in this study. Comparisons of the modeling results in the finest domain \(d03\) with the](#)  
159 [hourly observation data are performed in Shanghai, Hangzhou and Nanjing, and Hangzhou for](#)  
160 [2-m air temperature, 2-m relative humidity, 10-m wind speed, 10-m wind direction, surface O<sub>3</sub> and](#)  
161 [surface NO<sub>2</sub>. As a supplement, the modeling results and observations for the surface hourly O<sub>3</sub>](#)  
162 [concentrations in Wuxi are compared as well due to the fact that the meteorological data of](#)  
163 [Wuxi is unavailable from the University of Wyoming. Shanghai \(31.40°N, 121.46°E\) is the](#)  
164 [most populous city in China and Asia, as well as and also a global financial and transportation](#)  
165 [center. Locating to the northwest of Shanghai, Nanjing \(32.00°N, 118.80°E\) is the capital of](#)  
166 [Jiangsu Province and the second largest commercial center in East China. Hangzhou \(30.23°N,](#)  
167 [120.16°E\) is the capital of Zhejiang Province and located to the southwest of Shanghai. These](#)  
168 [cities are the provincial capitals and the typical metropolis in the YRD region. They are highly](#)  
169 [urbanized and industrialized, and all suffer from severe O<sub>3</sub> pollution. Additionally, Wuxi \(31.62°N,](#)  
170 [120.27°E\) is located between Shanghai and Nanjing and close to the Taihu Lake.](#)

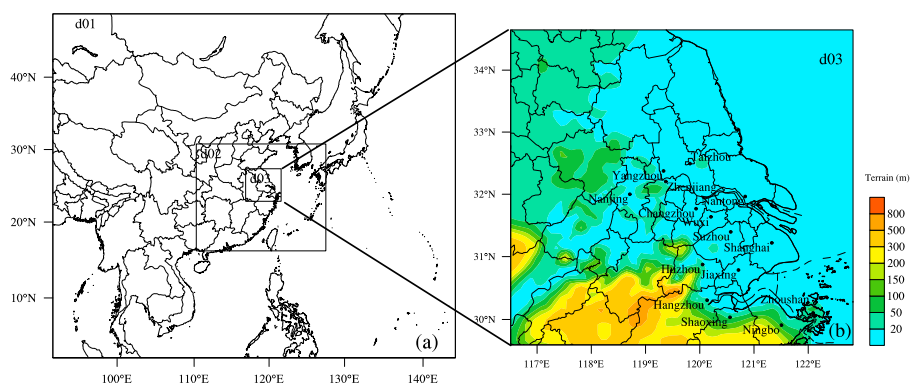
## 172 2.2 Model description and configurations

173 [In this paper study, WRF/CMAQ, which consists of the Weather Research and Forecasting](#)  
174 [\(WRF\) model version 3.4.1 and the Community Multi-scale Air Quality \(CMAQ\) Model version](#)  
175 [4.7.1, is applied to simulate the high O<sub>3</sub> episode over the YRD region in August 2013. Developed](#)  
176 [at the National Center for Atmospheric Research \(NCAR\), WRF is a new generation of](#)  
177 [meso-scale weather forecast model and assimilation system developed at the National Center for](#)  
178 [Atmospheric Research \(NCAR\). Numerous applications have proven that it shows a good](#)  
179 [performance in all kinds of weather forecasts and has broad application prospects in China \(Jiang](#)

180 et al., 2008; 2012; Wang et al., 2009b; Liu et al., 2013; Xie et al., 2014; 2016a; Liao et al., 2014;  
181 2015; Li et al., 2016; Zhu et al., 2016). WRF provides off-line meteorological fields as the input  
182 for the chemical transport model CMAQ. ~~The CMAQ modeling system~~ is a third generation of  
183 regional air quality model developed by the Environmental Protection Agency of USA (USEPA).  
184 A set of up-to-date compatible modules and control equations for the atmosphere is incorporated  
185 in the model, which can fully consider atmospheric complicated physical processes, chemical  
186 processes and the relative contribution of different species (Byun and Schere, 2006; Foley et al.,  
187 2010). CMAQ has been widely applied in China and proven to be a reliable tool in simulating air  
188 quality from city scale to meso scale (Li et al., 2012; Wei et al., 2012; Liu et al., 2013; Zhu et al.,  
189 2016).

190 The simulation run is conducted from 08:00 ([local standard time](#), LST) on August 2~~nd~~ to  
191 08:00 (LST) on August 16~~th~~ 2013, in which the first 48 h is taken as the spin-up time. Three  
192 one-way nested domains are used in WRF with a Lambert Conformal map projection. The domain  
193 setting is shown in Fig. 1. The outermost domain (domain 1, d01) covers the most areas of East  
194 Asia and South Asia, with the horizontal grids of 88×75 and the grid spacing of 81km. The nested  
195 domain d02 covers the southeastern part of China, with the horizontal grids of 85×70 and the grid  
196 spacing of 27km. The finest domain (domain 3, d03) covers the core areas of the YRD region,  
197 with the grid system of 70×64 and the resolution of 9km. For all domains, there are 23 vertical  
198 sigma layers from the surface to the top pressure of 100hPa, with about 10 layers in the [planetary](#)  
199 [PBL boundary layer](#). The detailed configuration options for the dynamic parameterization in WRF  
200 are summarized in Table 1. Additionally, the SLAB scheme that does not consider urban canopy  
201 parameters is adopted to model the urban effect. In order to reflect the rapid urban expansion in  
202 the YRD region, the default [United States Geological Survey \(USGS\)](#) land-use archives are  
203 updated by adding the present urban land-use conditions from 500-m Moderate Resolution  
204 Imaging Spectroradiometer (MODIS) data, based on the work of Liao et al. (2014; 2015). The  
205 initial meteorological fields and boundary conditions are from [1° resolution global reanalysis data](#)  
206 ~~provided by National Center for Environmental Prediction (NCEP) FNL global reanalysis data~~  
207 ~~with 1°×1° resolution~~. The boundary conditions are forced every 6 h.

208



209  
 210 **Fig. 1. Domain settings, include (a) the three nested modeling domains and (b) the nested domain 3 (d03)**  
 211 **with the terrain elevations and the locations of 15 main cities in the YRD region.**

212

213 **Table 1. The grid settings and the physical options for WRF in this study.**

Items	Options
Dimensions (x, y)	(88, 75), (85, 70), (70, 64)
Grid spacing (km)	81, 27, 9
Microphysics	WRF Single-Moment 5-class scheme (Hong et al., 2004)
Longwave Radiation	RRTM scheme (Mlawer et al., 1997)
Shortwave Radiation	Goddard scheme (Kim and Wang, 2011)
Surface layer	Moni-Obukhov scheme (Monin and Obukhov, 1954)
Land-surface layer	Noah Land Surface Model (Chen and Dudhia, 2001)
Planetary Boundary layer	YSU scheme (Hong et al., 2006)
Cumulus Parameterization	Grell-Devenyi ensemble scheme (Grell and Devenyi, 2002)

214

215 With respect to the air quality model, CMAQ uses the same vertical levels and the similar  
 216 three nested domains as those adopted in the meteorological simulation, whereas the CMAQ  
 217 domains are one grid smaller than the WRF domains. The Meteorology Chemistry Interface  
 218 Processor (MCIP) is used to convert WRF outputs to the input meteorological files needed by  
 219 CMAQ. The Carbon Bond 05 chemical mechanism (CB05) (Yarwood et al., 2005) is chosen for  
 220 gas-phase chemistry, and the 4rd generation CMAQ aerosol module (Byun and Schere, 2006) is  
 221 adopted for aerosol chemistry. The initial and outmost boundary conditions are obtained from the  
 222 Model for Ozone and Related Chemical Tracers version 4 (MOZART-4) (Emmons et al., 2010),  
 223 while those for the two nested inner domains are extracted from the immediate concentration files  
 224 of their parent domains. The anthropogenic emissions are mainly from the 2012-year



225 Multi-resolution Emission Inventory for China (MEIC) with  $0.25^{\circ} \times 0.25^{\circ}$  resolution, which is  
226 re-projected for the grids of China in both domains. For the grids outside of China, the inventory  
227 developed for the Intercontinental Chemical Transport Experiment-Phase B (INTEX-B) by Zhang  
228 et al. (2009) is used. The natural  $O_3$  precursor emissions are calculated by the natural emission  
229 model developed by Xie et al. (2007; 2009; 2014), including NO from soil, VOCs from  
230 vegetations, and  $CH_4$  from rice paddies and terrestrial plants. The biomass burning emissions are  
231 acquired from the work of Xie et al. (2014; 2016a).

### 232 **2.3 Integrated Process Rate (IPR) analysis method**

233 The CMAQ modeling system contains process analysis module (PROCAN), which consists  
234 of the Integrated Process Rate (IPR) analysis and the Integrated Reaction Rate (IRR) analysis  
235 (Byun and Schere, 2006). IPR has the capability of calculating the hourly contributions of  
236 individual physical processes and the net effect of chemical reaction compared to the overall  
237 concentrations, and thereby can determine the quantitative contribution of each process in a  
238 specific grid cell. The atmospheric processes ~~taken into consideration~~ considered in IPR include  
239 the horizontal advection (HADV), the vertical advection (ZADV), the horizontal diffusion (HDIF),  
240 the vertical diffusion (VDIF), the emissions (EMIS), the dry deposition (DDEP), the cloud  
241 processes with the aqueous chemistry (CLDS), the aerosol processes (AERO) and the gas-phase  
242 chemistry (CHEM). ~~The IPR analysis~~ has been widely applied to investigate the regional  
243 photochemical pollutions, and proven to be an effective tool to show the relative importance of  
244 every process and provide a fundamental interpretation (Goncalves et al., 2009; Li et al., 2012; Liu  
245 et al., 2013; Zhu et al., 2016).

246 In this paper, the period ~~during from~~ August 4- ~~to~~ 15 is selected for the IPR analysis. With  
247 the aid of IPR, we assess the roles of the individual physical and chemical processes involved in  
248  $O_3$  formation over the YRD region, and further present those in the typical cities ~~such as~~  
249 (Shanghai, Nanjing and Hangzhou). ~~Shanghai is the most populous city in China and Asia, as well~~  
250 ~~as a global financial and transportation center. Locating to the northwest of Shanghai, Nanjing is~~  
251 ~~the capital of Jiangsu Province and the second largest commercial center in East China. Hangzhou~~  
252 ~~is the capital of Zhejiang Province and located to the southwest of Shanghai. These cities are all~~  
253 ~~highly urbanized and industrialized, and suffer from severe  $O_3$  pollution.~~

### 254 **2.4 Evaluation method**

255 ~~Meteorological and air quality observation data are used to validate the reliability of~~  
 256 ~~simulation in this study.~~ Comparisons of the modeling results in the finest domain (d03) with the  
 257 hourly observation data are performed ~~in Shanghai (31.40°N, 121.46°E), Hangzhou (30.23°N,~~  
 258 ~~120.16°E) and Nanjing (32.00°N, 118.80°E) for 2-m air temperature, 2-m relative humidity,~~  
 259 ~~surface O<sub>3</sub> and surface NO<sub>2</sub> meteorological factors and air pollutants in Shanghai, Hangzhou, and~~  
 260 ~~Nanjing. As a supplement, Additionally, the modeling results and observations for the surface~~  
 261 ~~hourly O<sub>3</sub> concentrations in Wuxi (31.62°N, 120.27°E) are compared as well due to the fact that~~  
 262 ~~the meteorological data of Wuxi is unavailable from the University of Wyoming.~~ The  
 263 correlation coefficient (R), the normalized mean bias (NMB) and the root-mean-square error  
 264 (RMSE) are used to evaluate the model performance. These statistic values are calculated as  
 265 follows:

$$266 \quad R = \frac{\sum_{i=1}^N (S_i - \bar{S})(O_i - \bar{O})}{\sqrt{\sum_{i=1}^N (S_i - \bar{S})^2} \sqrt{\sum_{i=1}^N (O_i - \bar{O})^2}} \quad (1)$$

$$267 \quad NMB = \frac{\sum_{i=1}^N (S_i - O_i)}{\sum_{i=1}^N O_i} \times 100\% \quad (2)$$

$$268 \quad RMSE = \left[ \frac{1}{N} \sum_{i=1}^N (S_i - O_i)^2 \right]^{\frac{1}{2}} \quad (3)$$

269 ~~where~~  $S_i$  ~~represents the simulated value and~~  $O_i$  ~~represents the simulated and the~~ observed value,  
 270 ~~respectively.~~  $N$  means the total number of valid data. Generally, the model performance is  
 271 acceptable if the values of NMB and RMSE are close to 0 and that of R is close to 1.

272

### 273 3. Characteristics of the continuous ozone episode

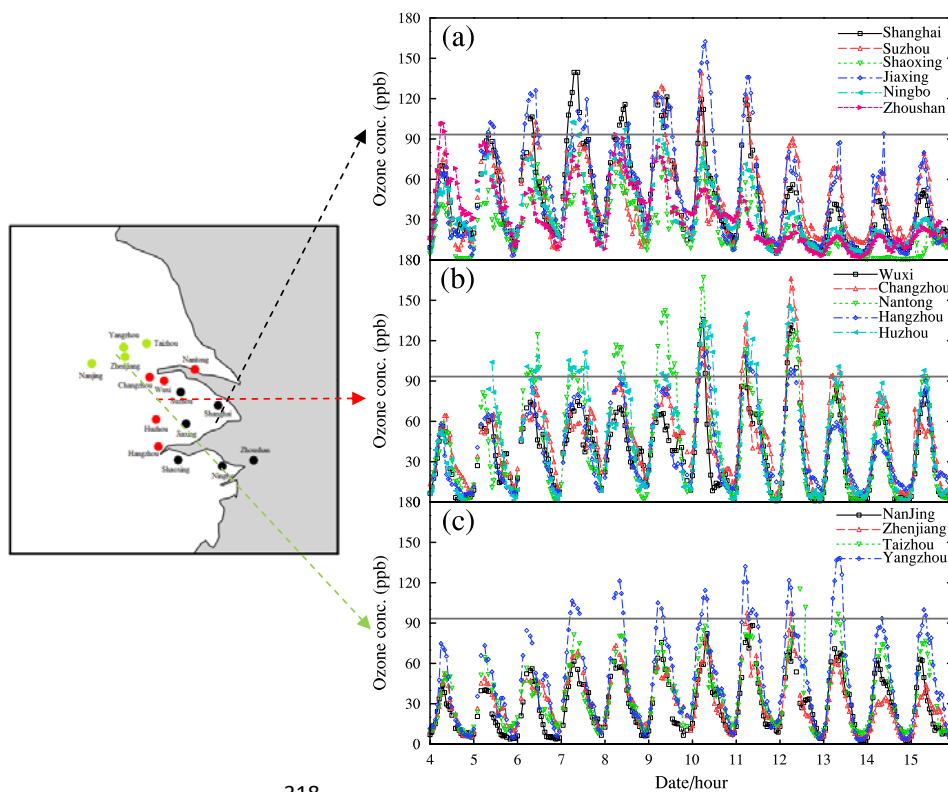
#### 274 3.1 Basic characteristic of the regional ozone episode in August 2013

275 Fig. 2 shows the temporal variation of the hourly O<sub>3</sub> concentrations observed ~~by in situ~~  
 276 ~~monitoring~~ in 15 typical cities over the YRD region from 00:00 (~~Universal Time Coordinated~~  
 277 UTC) ~~4~~ August ~~4~~ to 23:00 (UTC) ~~15~~ August ~~15~~ in 2013. Obviously, from August 7 to ~~August~~ 12,  
 278 high O<sub>3</sub> concentrations over 93.5 ppb (approximately equal to the ~~hourly~~ national air quality  
 279 standard of 200 µg/m<sup>3</sup> ~~for the hourly O<sub>3</sub> concentration~~) have been frequently recorded in 13 cities,  
 280 which means ~~O<sub>3</sub> concentrations in~~ most cities over the YRD region exceed the national air quality

281 standard. So, this high O<sub>3</sub> pollution episode is a typical regional O<sub>3</sub> pollution episode that can  
282 affect the people and the ecosystem in a large area. In general, for each city, there is a remarkable  
283 continuous growth in O<sub>3</sub> concentrations before the O<sub>3</sub> episode, followed by the lasting heavy O<sub>3</sub>  
284 pollution period. Though the O<sub>3</sub> concentrations in Shaoxing and Nanjing meet the national O<sub>3</sub>  
285 standard, their time series still show the similar tendency for the other cities in the same region.  
286 The excessive level of O<sub>3</sub> occurring in Huzhou, Jiaxing, Nantong, Yangzhou and Shanghai lasts for  
287 more than six consecutive days, reflecting the regional continuous characteristics of this O<sub>3</sub>  
288 pollution episode.

289 According to the temporal variation characteristics of O<sub>3</sub> illustrated in Fig. 2, the  
290 abovementioned 15 typical YRD cities can be classified into three categories: (1) the cities in the  
291 Southeast Coastal Region (SCR), including Shanghai, Suzhou, Jiaxing, Ningbo, Shaoxing, and  
292 Zhoushan; (2) the cities in the Central Inland Region (CIR), including Hangzhou, Huzhou, Wuxi,  
293 Changzhou, and Nantong; and (3) the cities in the Northwestern Inland Region (NIR), including  
294 Nanjing, Yangzhou, Zhenjiang, and Taizhou. The classification is primarily on basis of the  
295 observational facts that the maximum O<sub>3</sub> concentrations occur on August 10-11, 12, and 13, and  
296 begin to synchronously decrease on August 12, 13 and 14 in SCR, CIR and NIR, respectively. As  
297 shown in Fig. 2, in the Southeast Coastal Region (SCR), Zhoushan firstly exceeds the national O<sub>3</sub>  
298 standard on August 4, followed by Jiaxing, Shanghai, Suzhou and Ningbo. The peak hourly O<sub>3</sub>  
299 concentration of SCR occurs in Jiaxing on August 10, with the value up to 162.4 ppb. In the  
300 Central Inland Region (CIR), Huzhou is the first city exceeding the national O<sub>3</sub> standard, followed  
301 by the order of Nantong, Changzhou, Wuxi, and Hangzhou. The high-level O<sub>3</sub> pollution in  
302 Huzhou lasts during August 5-13. In Nantong and Changzhou, the maximum hourly O<sub>3</sub>  
303 concentration reaches 167.1 ppb on August 10 and 166.1 ppb on August 12, respectively. As for  
304 the Northwest Inland Region (NIR), Yangzhou, Zhenjiang, and Taizhou successively exceed the  
305 national O<sub>3</sub> standard. It is also noteworthy that the date when O<sub>3</sub> concentration exceeds the  
306 national air quality standard in coastal region is ahead of that in inland regions, so is the date of O<sub>3</sub>  
307 decrease. The different start time of O<sub>3</sub> decreasing in different regions might be related to the  
308 strong southeast wind in accordance with the movement of Typhoon Utor, which is discussed in  
309 Sect. 3.2 in detail. Table 2 presents the highest and the average concentrations of O<sub>3</sub>, as well as its  
310 precursors (NO<sub>2</sub> and CO), observed in these 15 cities during August 7-12 2013. The highest hourly

311  $O_3$  concentration occurs in Nantong with the value of 167.1 ppb, which is nearly 2 times of the  
 312 national air quality standard, followed by 166.1 and 162.4 ppb in Changzhou and Jiaxing,  
 313 respectively. It seems that  $O_3$  concentrations are higher in the cities around Shanghai, where the  
 314 concentrations of  $O_3$  precursors (shown in Table 2) and the water vapor are more adequate as well.  
 315 High concentrations of  $O_3$  and its precursors imply that there may be stronger photochemical  
 316 reactions in these cities.  
 317



318

319 Fig. 2. The time series of the observed  $O_3$  concentrations in 15 typical cities from August 4 to 15 August-2013  
 320 over the YRD region, which can be divided into three areas: (a) the Southeast Coast Region (SCR),  
 321 including Shanghai, Suzhou, Shaoxing, Jiaxing, Ningbo, and Zhoushan; (b) the Central Inland Region  
 322 (CIR), including Wuxi, Changzhou, Nantong, Hangzhou, and Huzhou; (c) the Northwest Inland Region  
 323 (NIR), including Nanjing, Zhenjiang, Taizhou and Yangzhou. The gray solid lines in (a), (b), and (c)  
 324 represent the national standard for the hourly  $O_3$  concentration, which is  $200 \mu\text{g}/\text{m}^3$ .  
 325

326 Table 2 presents the maximum and the average concentrations of  $O_3$  and  $\text{NO}_2$  in 15 YRD  
 327 cities during August 7-12 2013. It illustrates that the mean concentrations of  $\text{NO}_2$  in different YRD  
 328 cities range from 7.7 to 24.5 ppb during the  $O_3$  episode, indicating the heterogeneity of the spatial

带格式的: 下标

329 [distribution of O<sub>3</sub> precursor emissions. For O<sub>3</sub>, the highest hourly concentration \(167.1 ppb\)](#)  
 330 [occurs in Nantong, followed by 166.1 ppb in Changzhou and 162.4 ppb in Jiaxing. These values](#)  
 331 [are all nearly 2 times of the national air quality standard. It seems that O<sub>3</sub> concentrations are](#)  
 332 [higher in the cities around Shanghai, where the concentrations of O<sub>3</sub> precursors are more adequate](#)  
 333 [as well. High concentrations of O<sub>3</sub> and its precursors imply that there may be stronger](#)  
 334 [photochemical reactions.](#)

335  
 336 **Table 2. The maximum and average concentrations of O<sub>3</sub>, NO<sub>2</sub>, and CO observed in 15 cities during**  
 337 **August 7-12 2013 (ppb).**

Sites	O <sub>3</sub>		NO <sub>2</sub>		
	Max	Mean	Max	Mean	
Shanghai	139.5	55.1	35.1	15.6	
Southeast	Suzhou	139.1	50.9	50.6	19.7
Coast	Jiaxing	162.4	61.1	52.1	17.1
Region	Ningbo	113.4	41.9	31.2	12.4
(CSR)	Shaoxing	82.6	31.9	27.8	12.7
	Zhoushan	93.6	35.5	27.3	7.8
Central	Hangzhou	111.5	48.6	30.2	16.7
Inland	Huzhou	145.6	57.2	43.8	20.8
Region	Wuxi	135.8	43.2	39.9	18.8
(CIR)	Changzhou	166.1	55.7	58.4	24.5
	Nantong	167.1	56.0	48.2	20.9
Northwest	Nanjing	88.2	34.1	41.4	21.9
Inland	Yangzhou	132.1	54.1	36.0	17.1
Region	Zhenjiang	97.5	37.7	38.5	20.1
(NIR)	Taizhou	115.3	40.5	18.5	7.7

338  
 339 ~~According to the temporal variation characteristics of O<sub>3</sub> illustrated in Fig. 2, the~~  
 340 ~~abovementioned 15 typical YRD cities can be classified into three categories: (1) the cities in the~~  
 341 ~~Southeast Coastal Region (SCR), including Shanghai, Suzhou, Jiaxing, Ningbo, Shaoxing, and~~  
 342 ~~Zhoushan; (2) the cities in the Central Inland Region (CIR), including Hangzhou, Huzhou, Wuxi,~~  
 343 ~~Changzhou, and Nantong; and (3) the cities in the Northwestern Inland Region (NIR), including~~  
 344 ~~Nanjing, Yangzhou, Zhenjiang, and Taizhou. The classification is primarily on basis of the~~  
 345 ~~observational facts that the maximum O<sub>3</sub> concentrations occur on August 10, 11, 12, and 13, and~~

带格式表格

346 begin to synchronously decrease on August 12, 13 and 14 in SCR, CIR and NIR, respectively. As  
347 shown in Fig. 2, in the Southeast Coastal Region (SCR), Zhoushan firstly exceeds the national O<sub>3</sub>  
348 standard on August 4th, followed by Jiaxing, Shanghai, Suzhou and Ningbo. The peak hourly O<sub>3</sub>  
349 concentration of SCR occurs in Jiaxing on August 10 with the value up to 162.4 ppb. In the  
350 Central Inland Region (CIR), Huzhou is the first city exceeding the national O<sub>3</sub>-standard, followed  
351 by the order of Nantong, Changzhou, Wuxi and Hangzhou. The high level O<sub>3</sub>-pollution in Huzhou  
352 lasts from August 5th to 13th. In Nantong and Changzhou, the maximum hourly O<sub>3</sub>-concentration  
353 reaches 167.1 ppb on August 10 and 166.1 ppb on August 12, respectively. As for the Northwest  
354 Inland Region (NIR), Yangzhou, Zhenjiang and Taizhou successively exceed the national O<sub>3</sub>  
355 standard. It is also noteworthy that the date when O<sub>3</sub>-concentration exceeds the national air quality  
356 standard in coastal region is ahead of that in inland regions, so is the date of O<sub>3</sub>-decrease. The  
357 different time of the O<sub>3</sub> decrease in different regions might be related to the strong southeast  
358 wind in accordance with the movement of Typhoon Utor, which is discussed in Sect. 3.2 in detail.

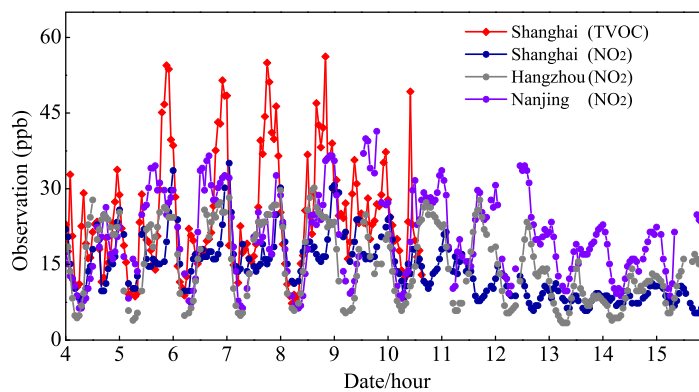
359 In general, for each city, there is a remarkable continuous growth in O<sub>3</sub> concentrations before  
360 the O<sub>3</sub>-episode, followed by the lasting heavy O<sub>3</sub> pollution period. Though the O<sub>3</sub>-concentrations in  
361 Shaoxing and Nanjing meet the national O<sub>3</sub> standard, their time series still show the similar  
362 tendency for the other cities in the same region. The excessive level of O<sub>3</sub> occurring in Huzhou,  
363 Jiaxing, Nantong, Yangzhou and Shanghai lasts for more than six consecutive days, reflecting the  
364 regional continuous characteristics of this O<sub>3</sub> pollution episode.

365 As for O<sub>3</sub> precursors, NO<sub>2</sub> and CO, their the averaged NO<sub>2</sub> concentrations in the YRD  
366 region during the O<sub>3</sub> episode show the variation range of approximately 7.7-24.5 and  
367 460.0-1094.0 ppb, respectively (Table 2), indicating the heterogeneity of the spatial emission  
368 distribution of O<sub>3</sub> precursors. Besides, the relative high hourly concentrations of NO<sub>2</sub> show good  
369 agreements with those of O<sub>3</sub>, implying it is one of the important O<sub>3</sub> precursor. Furthermore,  
370 Fig. 3 demonstrates the hourly variations of the observed NO<sub>2</sub> concentrations in Shanghai,  
371 Nanjing and Hangzhou from August 4 to 15 2013, and the time series of TVOC observed at SAES  
372 in Shanghai from August 4 to 10 2013. Obviously, there are two peaks in the diurnal cycles of  
373 NO<sub>2</sub> and VOC at all sites, which should be related with the rush hours in cities. The photolysis of  
374 NO<sub>2</sub> dominates O<sub>3</sub>-VOC-NO<sub>x</sub> chemistry after 8:00, and thereby makes the concentrations of  
375 precursors (NO<sub>2</sub> and VOC) begin to decrease. Thus, the related reactions form O<sub>3</sub> and increase its

带格式的：正文，左，行距：1.5  
倍行距，制表位：不在 35.16 字  
符

376 concentration until about 14:00. These diurnal variations of  $O_3$  and its precursors follow the  
377 typical patterns in the polluted areas and reflect the close relationships between  $O_3$ , VOC and  $NO_x$   
378 (Wang et al., 2013; Xie et al., 2016b). Moreover, the daily variations of  $NO_2$  and VOC show good  
379 agreement with those of  $O_3$ . For VOC, the concentration in Shanghai largely increases since  
380 August 6, which corresponds well with the over-standard  $O_3$  concentrations since then (Fig. 2).  
381 For  $NO_2$ , the higher values occur from August 6 to 11 in all cities, but the concentrations start to  
382 decrease on August 12, 13 and 14 in Shanghai, Hangzhou and Nanjing, respectively. It seems that  
383 the changes of  $O_3$  precursors ( $NO_2$  and VOC) are also affected by the movement of Typhoon Utor.

384  
385



386  
387 **Fig. 3. Temporal variations of the observed  $NO_2$  concentrations at Shanghai, Nanjing and Hangzhou**  
388 **stations from August 4 to 15 2013 and the observed TVOC concentration at SAES (31.17°N, 121.43°E) in**  
389 **Shanghai station from August 4 to 16:00 (UTC) August-10 2013.**

390

### 391 3.2 Meteorological condition and its effect

392 Favorable weather conditions have large impacts on the formation of severe  $O_3$  pollution  
393 (Huang et al., 2005; 2006; Wang et al, 2006b; Jiang et al., 2008; Cheng et al., 2014; Hung and Lo,  
394 2015). High-level  $O_3$  episodes often take place in hot seasons, when the meteorological conditions  
395 with high temperature and strong solar radiation are beneficial to the photochemical reactions of  
396  $O_3$  (Lam et al., 2005). Fig. 4 shows the variations of the surface meteorological parameters that  
397 are related to this photochemical pollution episode during August 4-15, including 2-m air  
398 temperature, 2-m relative humidity, 10-m wind speed and 10-m wind direction at the

带格式的: 字体: Times New Roman, 下标

带格式的: 字体: Times New Roman, 下标

带格式的: 字体: Times New Roman, 下标

带格式的: 字体: Times New Roman, 下标

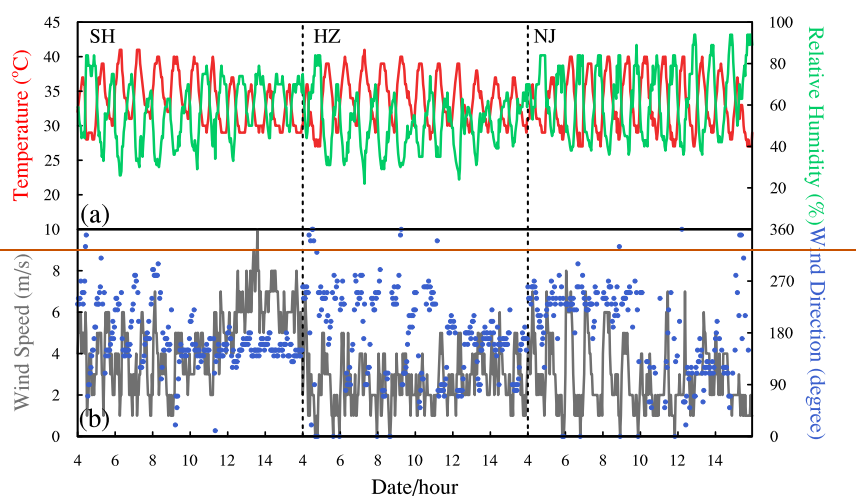
带格式的: 正文, 左, 制表位: 不在 35.16 字符

399 meteorological sites ~~of SH (31.40°N, 121.46°E) located~~ in Shanghai (SH) of SCR, ~~HZ (30.23°N,~~  
400 ~~120.16°E) located in~~ Hangzhou (HZ) of CIR, and ~~NJ (32.00°N, 118.80°E) located in~~ Nanjing (NJ)  
401 of NIR .

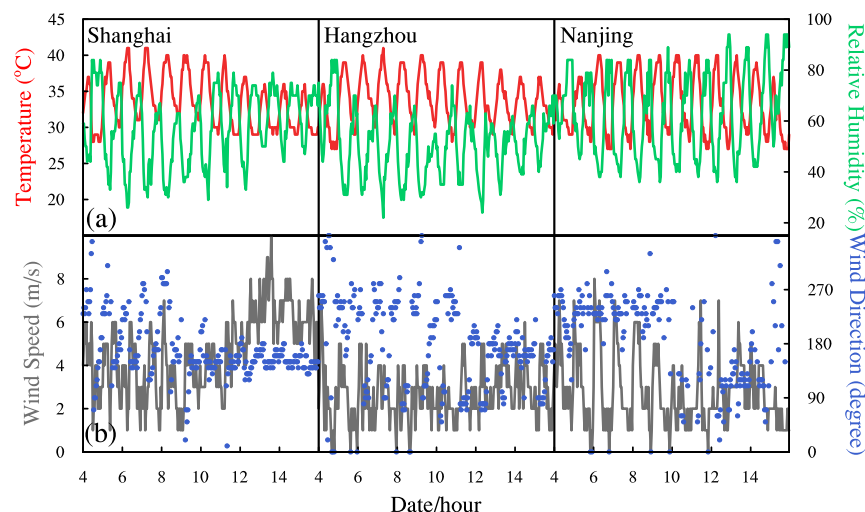
402 As shown in Fig. 43a, the hot weather at SH, HZ and NJ exists for nearly a week from  
403 August 7 to 12, with the hourly maximum temperature reaching the value over 40 °C. ~~°C.~~  
404 Meanwhile, the variations of 2-m relative humidity show the negative correlation with those of  
405 2-m air temperature. The minimum 2-m relative humidity at SH and HZ occur on August 9 and  
406 August 10 respectively, with the value below 75%. These minimum values are also lower than the  
407 values before and after the O<sub>3</sub> episode, suggesting that high-level O<sub>3</sub> episodes usually occur under  
408 the weather conditions with high temperature and low humidity. The value of 2-m relative  
409 humidity at NJ is relatively higher than those at SH and HZ and remains more stable. This  
410 extremely hot and dry weather condition at SH, HZ, and NJ are successively relieved on August  
411 12, 13 and 15, which coincide well with the reduction of surface O<sub>3</sub> concentrations in Shanghai,  
412 Hangzhou, and Nanjing (Fig. 2). With respect to the observed surface wind (Fig. 43b), the 10-m  
413 wind speed at SH, ~~and HZ, and NJ~~ is comparatively lower during the period of the O<sub>3</sub> episode,  
414 while it is suddenly intensified after August 12. Meanwhile, the wind direction is fluctuating from  
415 ~~August 7 to 12-August,~~ while it maintains southeasterly wind after August 12 as well. The growth  
416 of wind speed is more distinct at SH, with the maximum value of approximately 10 m/s. The wind  
417 speed at NJ has an obviously diurnal variation from August 4 to 8, and the minimum value occurs  
418 on August 10.

419





420



421

422 **Fig. 43.** Temporal variations of the main meteorological parameters at **ShanghaiH** ( $31.40^{\circ}\text{N}, 121.46^{\circ}\text{E}$ ),  
 423 **HangzhouZ** ( $30.23^{\circ}\text{N}, 120.16^{\circ}\text{E}$ ) and **NanjingJ** ( $32.00^{\circ}\text{N}, 118.80^{\circ}\text{E}$ ) meteorological stations during August  
 424 4-15, 2013, including (a) 2-m air temperature (the red solid line) and 2-m relative humidity (the green solid  
 425 line); (b) 10-m wind speed (the gray solid line) and 10-m wind direction (the blue scatter points).

426

427 Fig. 45 displays the weather charts for the 500hPa layer over the East Asia at 00:00 (UTC) on  
 428 August 6, 8, 10, and 12 August 2013, which can illustrate the main synoptic patterns causing the  
 429 O<sub>3</sub> pollution. Obviously, during the period of the selected O<sub>3</sub> episode, the whole YRD region is  
 430 under the control of the strong Western Pacific subtropical high, which is stronger and extends  
 431 much farther west than normal. The anomaly of the subtropical high might be the direct and

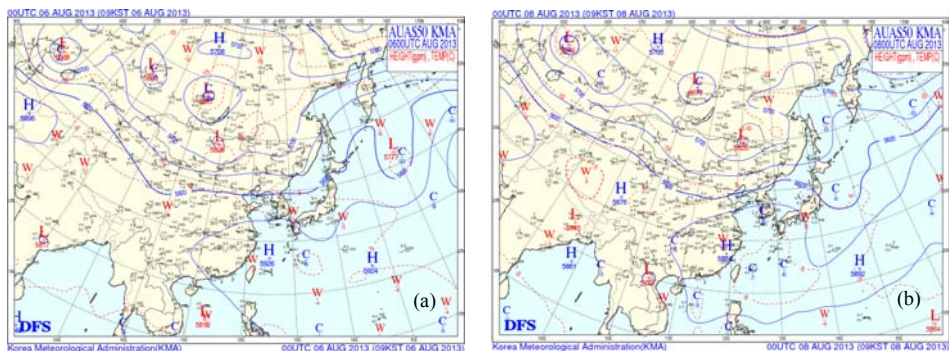
432 | leading cause of the abnormally high temperature shown in Fig. 43a (Peng et al., 2014). The  
433 | intensity of the subtropical high is usually characterized by the area index, defined as the total  
434 | number of grid points that have geopotential heights of 588 decameters or greater in the region of  
435 | 110-180°E and northward of 10°N. As shown in Fig. 54, the 588-decameter area covers most of  
436 | southeast China, and the high pressure center (592-decameter area) is located in the southeastern  
437 | coastal areas as well as the surrounding sea areas, which means the subtropical high is very  
438 | intensive. This high pressure strengthens and remains over the YRD region for several days (from  
439 | August 6 to 12), implying that the air subsides to the ground. The downward air acts as a dome  
440 | capping the atmosphere, and helps to trap heat as well as air pollutants at the surface. Without the  
441 | lift of air, there is little convection and therefore little cumulus clouds or rains. The end result is a  
442 | continual accumulating of solar radiation and heat on the ground, which may greatly enhance the  
443 | photochemical reactions between the abundant build-up air pollutants.

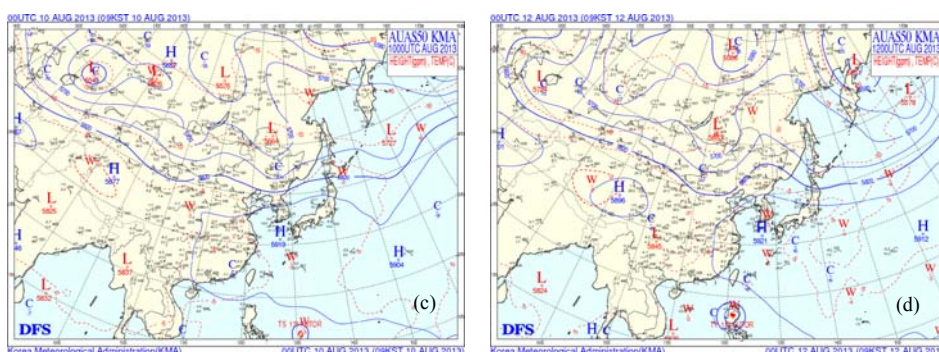
444 | The other weather system worthy of note is Typhoon Utor (shown in Fig. 54c and d).  
445 | Typhoon Utor is one of the strongest typhoons in the 2013 Pacific typhoon season, ~~with the~~  
446 | ~~international code of 1311~~. It is formed early on August 8, develops into a tropical storm on  
447 | August 9, undergoes an explosive intensification within a half of day, and achieves typhoon status  
448 | on early August 10. After landing in Luzon of the Philippines on late August 11, it reemerges in  
449 | the South China Sea on August 12. Typhoon Utor hits the land of Guangdong Province in  
450 | China on August 14, and thereby is finally weakened into a tropical storm. In the end, it is  
451 | ultimately dissipated on August 18. It was reported that ozone episodes during the hot season are  
452 | usually associated with the passage of tropical cyclones close to the territory (Huang et al., 2005;  
453 | Wang et al., 2006b; Jiang et al., 2008; Cheng et al., 2014; Hung and Lo, 2015). When a site is at  
454 | the front of moving typhoon system, it can be controlled by the downward airflow induced by the  
455 | typhoons' peripheral circulation. So, the typhoon system can cause the local weather around the  
456 | site with high temperature, low humidity, strong solar radiation and small wind for a short time,  
457 | before it is close enough to bring winds and rains. All these changes of meteorological conditions  
458 | can ~~significantly affect~~ ~~help to the formation of~~ ~~form the~~ severe continuous O<sub>3</sub> pollution (Jiang et al.,  
459 | 2008). In this O<sub>3</sub> episode, the YRD region may be influenced by the peripheral circulation of  
460 | Typhoon Utor as well. Especially on August 10-11, the downward airflow in the troposphere is  
461 | significantly strengthened (shown in Fig. 6-7 ~~and detailedly discussed in Sect. 4.2~~), which may

462 | enhance the build-up of heat and air pollutants, and thereby result in worse air pollution-quality  
463 | shown in Fig. 2.

464 | Moreover, from August 12 to 14 (shown in Fig. 4d5d), with the approaching of Typhoon Utor  
465 | from August 12 to 14, the near-surface breeze over the YRD region gradually turns to be the  
466 | prevailing southeasterly or southerly wind (Fig. 5d), with the highest wind speed up to 6-10 m/s in  
467 | Shanghai (Fig. 4). The strengthened wind can bring the clean marine air from ocean to inland, and  
468 | thereby effectively mitigate the O<sub>3</sub> pollution. Meantime, Typhoon Utor also gradually affects the  
469 | position and strength of the Western Pacific subtropical high. As the typhoon continuously  
470 | approaching and finally landing on Guangdong, the high pressure system is forced to retreat  
471 | easterly and move northwards. When the high pressure center completely moves to the oceans, the  
472 | YRD region is totally under the control of the typhoon system. In the end, the hot weather is  
473 | relieved and the O<sub>3</sub> pollution is mitigated. The coastal cities in CSR are closer to the typhoon  
474 | system, so they are firstly influenced during this period. Thus, the wind at SH in CSR firstly  
475 | changes, followed by HZ in CIR and NJ in NIR. In the same way, 2-m air temperature and O<sub>3</sub>  
476 | concentrations also successively decrease from southeast (SH in CSR) to northwest (NJ in NIR)  
477 | owing to the scavenging effect.

478





479 | Fig. 45. Weather charts at the 500hPa layer over the East Asia at 00:00 (UTC) on (a) August 6, (b) August 8,  
 480 | (c) August 10, and (d) August 12 2013 (from Korea Meteorological Administration).

481

## 482 | 4 Modeling results and discussions

### 483 | 4.1 Evaluation of model performance

484 | To evaluate the simulation performance, the hourly modeling results during the period of  
 485 | ~~August~~ 4-15 ~~August~~ 2013 are compared with the observation records. Table 3 presents the  
 486 | performance statistics, including the values of the correlation coefficient (R), the normalized mean  
 487 | bias (NMB), and the root-mean-square error (RMSE), which are all calculated for 2-m air  
 488 | temperature ( $T_2$ ), 2-m relative humidity ( $RH_2$ ), 10-m wind speed ( $Wspd_{10}$ ), 10-m wind direction  
 489 | ( $Wdir_{10}$ ), surface ozone concentrations ( $O_3$ ), and surface nitrogen dioxide concentrations ( $NO_2$ ) in  
 490 | Shanghai (SH), Nanjing (NJ), and Hangzhou (HZ).

491 | As indicated in Table 3, the simulated results of surface air temperature and relative humidity  
 492 | from WRF show good correlation agreement with the observations. The highest correlation  
 493 | coefficient of 2-m air temperature ( $T_2$ ) is found to be 0.91 at SH, followed by 0.84 at NJ and 0.80  
 494 | at HZ (statistically significant at 95% confident level). The corresponding correlation coefficients  
 495 | for 2-m relative humidity ( $RH_2$ ) are 0.85, 0.83 and 0.78, respectively. The values of RMSE for  $T_2$   
 496 | at SH, NJ and HZ are 4.15, 2.91 and 3.09 ~~°C~~ °C, and those for  $RH_2$  are 19.3%, 9.41% and 13.96%  
 497 | respectively. ~~However, our~~ simulation underestimates  $T_2$  and overestimates  $RH_2$  to some  
 498 | certain extent, with the values of NMB for  $T_2$  at SH, NJ and HZ being -5.68%, -5.98% and -6.53%,  
 499 | and those for  $RH_2$  being 12.64%, 4.52% and 16.36%. These biases might be attributed to the  
 500 | uncertainty caused by the SLAB scheme, which can underestimate temperature in summer (Liao  
 501 | et al., 2014). ~~NMB (1.53%, 5.92% and 9.21% at SH, NJ and HZ site, respectively) and RMSE~~  
 502 | (2.18, 2.41 and 2.39 at SH, NJ and HZ site, respectively) are displayed concerning the wind speed

带格式的: 字体: Times New Roman, 非上标/下标

503 ~~variable. The simulated wind speeds are a little higher than observed results, but still reasonable.~~  
504 ~~Owing to the vectorial nature of wind, statistical comparisons for wind direction are shown~~  
505 ~~excluding mean values of observation and simulation data, with the values of NMB at SH, NJ and~~  
506 ~~HZ site being 3.53%, -1.57% and -0.13%, respectively. However, According to the relevant~~  
507 studies (Li et al., 2012; Liao et al., 2015; Xie et al., 2016a), this level of over- or under-estimation  
508 is still acceptable. ~~The wind components are closely related to the transport processes. As shown~~  
509 ~~in Table 3, our modeling results of wind speed and direction basically reflect the characteristics of~~  
510 ~~wind fields. For 10-m wind speed ( $W_{spd10}$ ), R is 0.77 at SH, 0.74 at NJ, and 0.75 at HZ,~~  
511 ~~respectively. Though the values of NMB (1.53%, 5.92%, and 9.21%) and RMSE (2.18, 2.41 and~~  
512 ~~2.39) display that the simulated wind speeds are a little overestimated, the biases are still~~  
513 ~~reasonable and acceptable. For 10-m wind direction ( $W_{dir10}$ ), the simulated values also fit the~~  
514 ~~observation records well, with the R values of 0.63 at SH, 0.57 at NJ and 0.58 at HZ. Comparing~~  
515 ~~the mean values from SIM and OBS, we can find that WRF model generally simulates the~~  
516 ~~prevailing wind direction during this period.~~ In summary, the abovementioned performance  
517 statistics numbers ~~basically~~ illustrate that the WRF simulation can reflect the major characteristics  
518 of meteorological conditions ~~during of~~ this O<sub>3</sub> episode, and the meteorological outputs can be used  
519 in the pollutant concentration simulation.

520

521 **Table 3. Comparisons between the simulations and the observations at Shanghai, Nanjing and Hangzhou**  
522 **stations during August 4-15 2013.**

Sites <sup>a</sup>	Vars <sup>b</sup>	Mean		R <sup>c</sup>	NMB <sup>f</sup>	RMSE <sup>g</sup>
		OBS <sup>c</sup>	SIM <sup>d</sup>			
SH	T <sub>2</sub> (°C)	33.27	31.38	0.91	-5.68%	4.15
	RH <sub>2</sub> (%)	57.91	65.23	0.85	12.64%	19.3
	$W_{spd10}$ (m s <sup>-1</sup> )	4.59	4.66	0.77	1.53%	2.18
	$W_{dir10}$ (°)	176.34	182.57	0.63	3.53%	41.44
	O <sub>3</sub> (ppb)	87.77	82.5	0.81	-6.00%	38.79
	NO <sub>2</sub> (ppb)	29.01	38.25	0.54	31.85%	28.95
	T <sub>2</sub> (°C)	32.95	30.98	0.84	-5.98%	2.91
NJ	RH <sub>2</sub> (%)	63.28	66.14	0.83	4.52%	9.41
	$W_{spd10}$ (m s <sup>-1</sup> )	3.21	3.4	0.74	5.92%	2.41
	$W_{dir10}$ (°)	197.68	194.58	0.57	-1.57%	71.19
	O <sub>3</sub> (ppb)	69.7	78.15	0.81	12.12%	36.8
	NO <sub>2</sub> (ppb)	41.44	40.09	0.61	-3.26%	22.4

	T <sub>2</sub> (°C)	33.25	31.08	0.8	-6.53%	3.09
	RH <sub>2</sub> (%)	52.76	61.39	0.78	16.36%	13.96
HZ	Wspd <sub>10</sub> (m s <sup>-1</sup> )	3.04	3.32	0.75	9.21%	2.39
	Wdir <sub>10</sub> (°)	186.45	186.2	0.58	-0.13%	69.44
	O <sub>3</sub> (ppb)	76.57	84.51	0.83	10.37%	33.95
	NO <sub>2</sub> (ppb)	31.06	27.21	0.66	-12.40%	16.86

523 <sup>a</sup> Sites indicates the city where the observation sites locate, including Shanghai (SH), Nanjing (NJ), and Hangzhou  
524 (HZ); <sup>b</sup> Vars indicates the variables under validation, including 2-m air temperature (T<sub>2</sub>), 2-m relative humidity  
525 (RH<sub>2</sub>), 10-m wind speed (Wspd<sub>10</sub>), 10-m wind direction (Wdir<sub>10</sub>), ozone (O<sub>3</sub>), and nitrogen dioxide (NO<sub>2</sub>). The  
526 words between the parentheses behind variables indicate the unit; <sup>c</sup> OBS indicates the observation data; <sup>d</sup> SIM  
527 indicates the simulation results from WRF/CMAQ/Chem; <sup>e</sup> R indicates the correlation coefficients, with  
528 statistically significant at 95% confident level; <sup>f</sup> NMB indicates the normalized mean bias; <sup>g</sup> RMSE indicates the  
529 root-mean-square error.

530

531 Fig. 5-6 shows the comparisons between the modeling results from CMAQ and the observed  
532 hourly concentrations of O<sub>3</sub> in Shanghai, Nanjing, and Hangzhou and Wuxi during August 4-15  
533 August-2013. Obviously, the observations and the simulated results present reasonable agreement  
534 at each site, with the correlation coefficients of 0.81 to 0.83, NMB of -6% to 12.12%, RMSE of  
535 33.95 to 38.79 ppb. Moreover, the simulation also reproduces the diurnal variation of O<sub>3</sub>, which  
536 shows that the concentration reaches its maximum at around noon time and gradually decreases to  
537 its minimum after midnight. With respect to the O<sub>3</sub> precursor, comparisons of NO<sub>2</sub> concentrations  
538 between simulation results and observations show that the correlation coefficient at each city is  
539 about 0.6 (given in Table 3), which further prove that the process of O<sub>3</sub> formation is captured  
540 reasonable well over the YRD region and throughout the episode. However, CMAQ overestimates  
541 NO<sub>2</sub> and underestimates O<sub>3</sub> in Shanghai, while underestimates NO<sub>2</sub> and overestimates O<sub>3</sub> in  
542 Nanjing and Hangzhou. These biases of O<sub>3</sub> and NO<sub>2</sub> can should mainly be attributed to the  
543 uncertainties in emissions of O<sub>3</sub> precursors (NO<sub>x</sub> and VOC<sub>x</sub>) (Li et al., 2012; Liao et al., 2015; Xie  
544 et al., 2016). Because of the VOC-sensitive O<sub>3</sub> chemistry in the daytime and NO<sub>x</sub> titration at night  
545 in the YRD region (Xie et al., 2014), higher estimation of NO<sub>x</sub> emission in Shanghai may lead to  
546 higher NO<sub>2</sub> and lower O<sub>3</sub> predictions, while lower NO<sub>x</sub> estimations in Nanjing and Hangzhou may  
547 result in lower NO<sub>2</sub> and higher O<sub>3</sub> modeling results. Trelated with O<sub>3</sub> precursor emissions,  
548 meteorology, and observation deviation (Li et al., 2012). In addition, note that this work is based  
549 on the 2012 emission scenario, while the simulation is conducted for a certain period in 2013 with

带格式的：下标

带格式的：下标

带格式的：下标

带格式的：下标

带格式的：下标

带格式的：下标

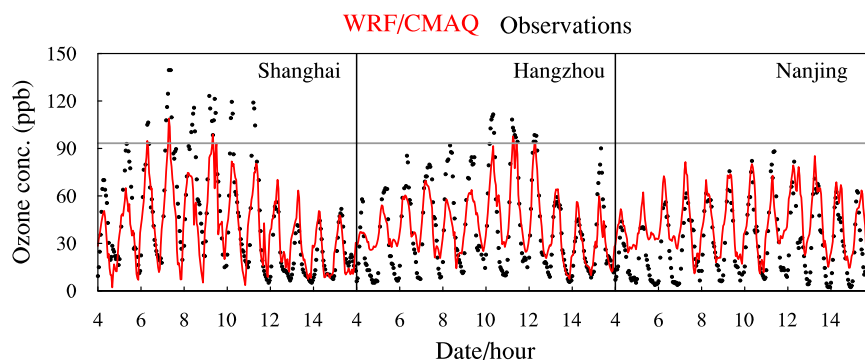
带格式的：下标

带格式的：下标

带格式的：下标

550 ~~the anthropogenic emission inventory not updated. The uncertainties in emissions of ozone~~  
 551 ~~precursors (NO<sub>x</sub> and VOCs) may cause these biases as well (Liao et al., 2015; Xie et al., 2016). he~~  
 552 ~~undervalued NO<sub>2</sub> and overvalued O<sub>3</sub> in Nanjing and Hangzhou. These underestimates should can~~  
 553 ~~also be related with the increasing overestimations in of WS<sub>10</sub> and the negative biases in T<sub>2</sub>.~~  
 554 Moreover, the uncertainties in nonlinear chemical reactions coupled in CMAQ may also have  
 555 important effects on model predictions. For example, the modeling results cannot catch the low O<sub>3</sub>  
 556 values observed at night in Nanjing (Fig. 64b) and Hangzhou (Fig. 64e) and Wuxi (Fig. 46d),  
 557 implying there may be some imperfections in the nocturnal chemistry of CMAQ. Nevertheless, the  
 558 performance of CMAQ model is comparable to the other applications (Goncalves et al., 2009; Li  
 559 et al., 2012; Zhu et al., 2016). Compared to these previous related studies, the simulation in this  
 560 study attains an acceptable and satisfactory result. Thus, the consistency of simulation and  
 561 observation demonstrates that the modeling results are capable of capturing and reproducing the  
 562 characteristics and changes of photochemical pollutants, and can be used to provide valuable  
 563 insights into the governing processes of this O<sub>3</sub> episode.

564

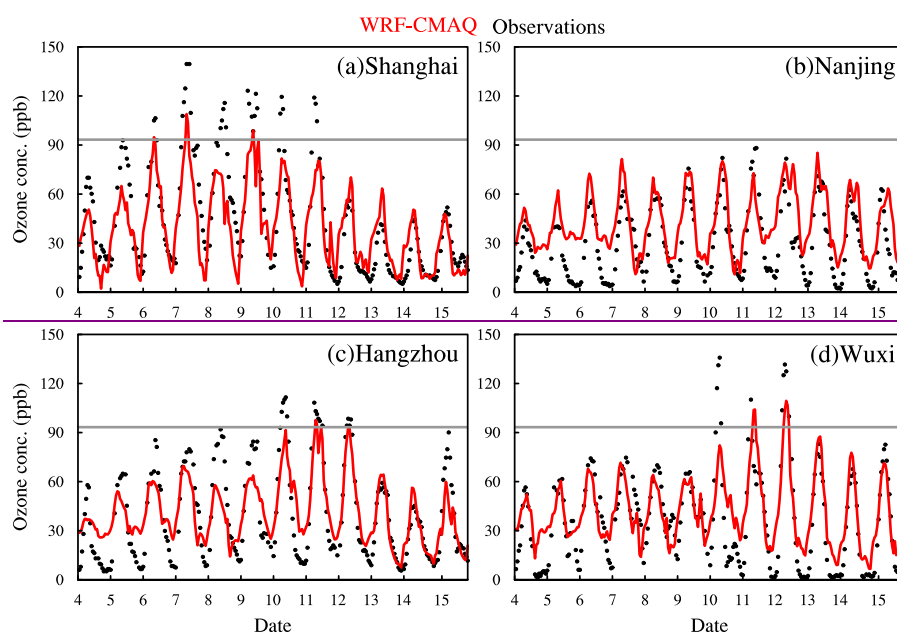


565

带格式的: 下标

带格式的: 下标





566

567 **Fig. 65.** Hourly variations of the observed and the simulated O<sub>3</sub> concentrations in (a) Shanghai (SH), (b)  
 568 Nanjing (NJ), (c) and Hangzhou (HZ), and (d) Wuxi. In (a), (b), (c), and (d), the red solid lines show the  
 569 modeling results, the black dot lines give the observations, and the solid gray lines represent the national  
 570 standard for the hourly O<sub>3</sub> concentration, which is 200 μg/m<sup>3</sup>.

571

#### 572 4.2 Characteristics of the vertical airflows

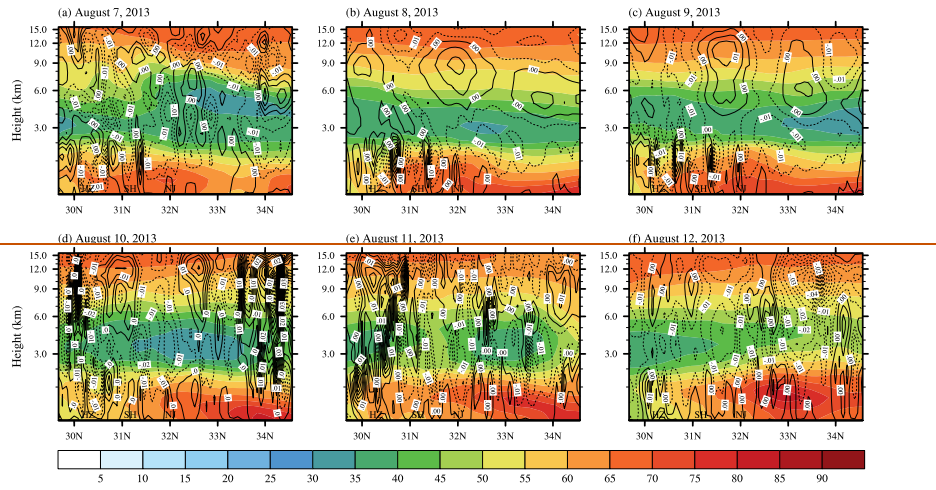
573 Fig. 67 presents the daytime vertical wind velocity as well as the vertical distribution of O<sub>3</sub>  
 574 concentrations from 29.67°NE to 34.76°EN during August 7-12 2013. Along the vertical  
 575 cross-section, the values from 118°E to 122°E are averaged in the meridional direction. Moreover,  
 576 Fig. 8 displays the daytime vertical wind velocity and vertical distribution of O<sub>3</sub> concentrations  
 577 from 116.5°E to 122.9°E in the zonal direction along the latitude of 31.40°N, (where Shanghai site  
 578 is located) lies during August 7-12 2013. The simulation results clearly illustrate that there are  
 579 strong downward airflows over the YRD region during the period of the regional high-level O<sub>3</sub>  
 580 pollution, which can be attributed to the fact that these areas are under the control of the  
 581 subtropical high and the sinking airflow is predominant (as discussed in Sect. 3.2).

582 From August 7 to 9 August 2013 (shown in Fig. 6a7a-c and 8a-c), except for the mentioned  
 583 regional sinking airflows, there are still some local thermal circulations, which are related with  
 584 urban heat islands, continually occurring at the lower atmospheric layers (< 2 km) along the

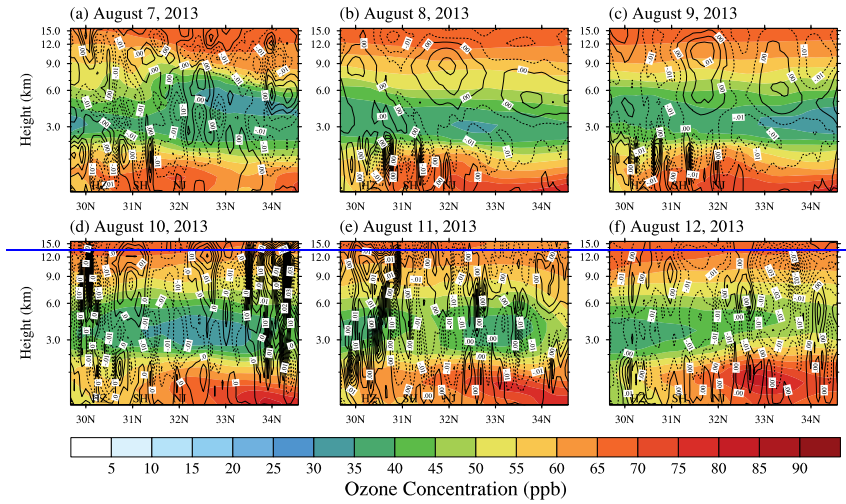


585 vertical cross-section of Hangzhou (HZ) — Shanghai (SH) - Nanjing (NJ). These circulations are  
586 related with urban heat islands. Usually high pressures are accompanied by more stagnant and fair  
587 dry weather, so the upward and the downward flows caused by urban-breeze circulations can  
588 easily appear in the urban areas ~~of SH, HZ, and NJ.~~ With respect to For the vertical distribution of  
589 O<sub>3</sub> in the meridional direction (Fig. 7a-e), its high concentrations (> 60-50 ppb) generally usually  
590 appear from the surface to 1.5km height above the cities, with the maximum values over 70 ppb in  
591 and around cities. While in the zonal direction of 31-40°N (Fig. 8a-e), high O<sub>3</sub> centers (> 90 ppb)  
592 appear near the surface in the southeast coastal region. As discussed in Sect. 3.2, induced by the  
593 regional sinking airflows, air pollutants ~~are~~ tend to be trapped on the ground due to the regional  
594 sinking airflows. Moreover, the local circulations over the cities (~~Fig. 6~~) make the urban areas to  
595 be the convergence zones, and thereby more air pollutants can be accumulated in and around these  
596 cities. Under the weather conditions induced by the subtropical high, such as high air temperature,  
597 stronger solar radiation and less water vapor, the chemical reactions between the build-up air  
598 pollutants can be enhanced to form the high-level O<sub>3</sub> pollution. Additionally, Fig. 7a-c also show  
599 that there are maximum O<sub>3</sub> concentrations (> 90 ppb) occurring near the surface in and around SH.  
600 This phenomenon should be explained by the fact that the coastal city (SH) is firstly affected by  
601 Typhoon Utor  
602 However, ~~f~~From August 10 to 12, with the approaching of Typhoon Utor, the vertical air  
603 movements over the YRD region are not restricted at the lower atmosphere any more. As shown in  
604 Fig. ~~7d-f and 8d-e f~~ (August 10 and 11), there are stronger downward airflows from the surface  
605 to the top of troposphere. As discussed in Sect. 3.2, ~~SH, HZ, and NJ~~ the YRD cities are at the front  
606 of the moving typhoon system, so the peripheral circulation of Typhoon Utor may enhance the  
607 sinking of atmosphere, which can lead to higher air temperature, lower humidity, and stronger  
608 solar radiation. Affected by the enhanced downward air movement as well as the relevant changes  
609 of meteorological conditions, O<sub>3</sub> concentrations over the YRD region maintain a high pollution  
610 level, with the O<sub>3</sub> concentrations over 60 ppb below the height of 1.5 km (Fig. 7d-f). Furthermore,  
611 ~~as shown in Fig. 8d to f,~~ the high value center of O<sub>3</sub> concentrations above (>90 ppb) moves ~~from~~  
612 ~~southeast to north~~westwards during August 10-12, implying that the peripheral circulation of  
613 Typhoon Utor can drive the air from the coastal areas to the inland areas.

带格式的：下标



615



616

617

618

619

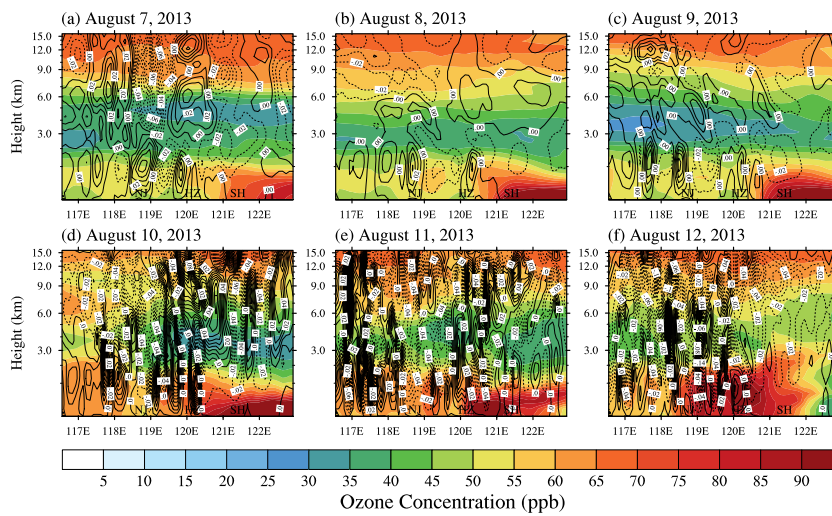
620

621

622

623

**Fig. 67. Simulated daytime vertical wind velocity and vertical distribution of O<sub>3</sub> concentrations in the YRD region from 29.67°E N to 34.76°E N during August 7 to 12 2013, with the values averaged in the meridional direction from 118°E to 122°E. The marks of HZ, SH, and NJ point out the latitudes of Hangzhou, Shanghai, and Nanjing, respectively. The solid lines show the zero velocity lines or positive wind speeds and represent upward airflow, while the dotted lines show the negative wind speeds and represent downward airflow. (anThe interval isof 0.005 m/s).**



带格式的：居中

624

625 **Fig. 78. Simulated daytime vertical wind velocity and vertical distribution of O<sub>3</sub> concentrations from**  
 626 **116.5°E to 122.9°E during August 7 to 12 2013 at the zonal direction along the latitude of 31.40°N; (where**  
 627 **Shanghai site lies located) during August 7 to 12 2013. The marks of SH-HZ, SHHZ and NJ point out the**  
 628 **longitudes of Shanghai-Hangzhou, HangzhouShanghai, and Nanjing, respectively. The dotted lines show the**  
 629 **negative wind speeds and represent downward airflow, while the solid lines show the zero-velocity lines**  
 630 **or the positive wind speeds and represent upward airflow zero vertical velocity, while the dotted lines show**  
 631 **the negative wind speeds and represent downward airflow. (The interval of is 0.01 m/s).**

632

633

634 The vertical changes of wind velocity and O<sub>3</sub> concentrations above Shanghai, Hangzhou and  
 635 Nanjing are further illustrated in Fig. 897. Similarly to that in Fig. 67 and 8, the atmospheric  
 636 subsidence can also be found in the boundary layer troposphere of Shanghai and Hangzhou (Fig.  
 637 97a and b) (usually occur at more than 1 km above the surface) during the period of the high-level  
 638 O<sub>3</sub> pollution (from August 7 to 12 August). For With respect to Shanghai site, affected by the  
 639 extremely high temperature, more active photochemical reactions lead to higher O<sub>3</sub> concentrations  
 640 in the whole atmospheric boundary layer. The downward airflows induced by the subtropical high  
 641 trap and enhance the accumulation of surface O<sub>3</sub> as time passes. Thus, high O<sub>3</sub> concentrations are  
 642 formed below 2 km above the urban areas of Shanghai, and the high concentration centers occur  
 643 near the surface below 500 m. It is interesting that O<sub>3</sub> concentration on August 8 is comparatively  
 644 lower, which can be seen in Fig. 2 as well. This phenomenon can be explained by the fact shown  
 645 in Fig. 97a that the transient upward airflow occurs at above 300 m over Shanghai and inhibits the  
 646 accumulation of the O<sub>3</sub> pollution at the surface (shown in Fig. 8a). Additionally, Fig. 897a also

647 presents the possible effects of Typhoon Utor on the formation of O<sub>3</sub>. On August 10, when the  
648 typhoon system approaches to the eastern coastal areas of China, the sinking air above Shanghai is  
649 apparently strengthened, and thereby enhances the intensity of O<sub>3</sub> pollution as well as the scope of  
650 the pollution. But after August 12, when Typhoon Utor changes the wind and even impacts the  
651 subtropical high, high temperature is alleviated and the build-up O<sub>3</sub> is transported to other places.

652 Thus, the pollution is mitigated. As to

653 As for Hangzhou site (Fig. 798b), from August 7 to 9, owing to weaker photochemical  
654 reactions, lower O<sub>3</sub> concentrations than that in Shanghai are found in the boundary layer owing to  
655 weaker photochemical reaction, while much more remarkable atmospheric subsidence is observed  
656 above Hangzhou, indicating that massive chemical formation is important for vertical transport  
657 process in the high O<sub>3</sub> episodes. However, the O<sub>3</sub> concentration can exceeds the national standard  
658 from August 10 to 12 (Fig. 2), which should be influenced by the typhoon system. The influence  
659 process is similar to the above discussion for Shanghai, that is, the upper downward airflows (over  
660 1 km above the surface) are enhanced significantly since August 10. The over standard of O<sub>3</sub>  
661 concentration can be attributed to the strengthened chemical formation, as well as the enhanced  
662 vertical transport process. The influence by the typhoon system is similar to the above discussion  
663 for Shanghai, as the downward air flow is enhanced significantly since August 10.

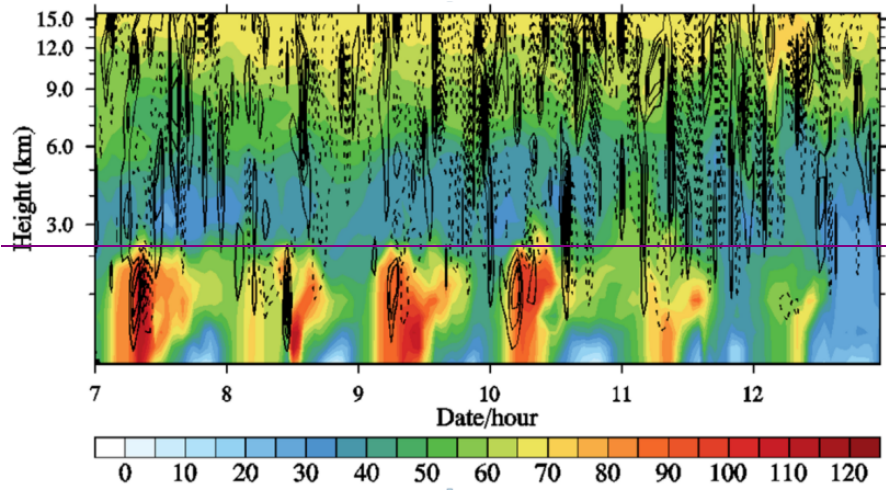
664 As demonstrated in Fig. 89e But for Nanjing, the O<sub>3</sub> concentration does not exceed the  
665 national O<sub>3</sub> standard during August 7-12 (Fig.2 and 8), which should be attributed to the fact that  
666 Nanjing is far away from the coastal areas and thereby hardly affected by the downward flow in  
667 the typhoon periphery. Though the O<sub>3</sub> concentration in Nanjing increases on August 12, it should  
668 mainly be caused by the local photochemical reactions because the vertical movement below 2 km  
669 above Nanjing is dominated by upward airflows.

670 the vertical O<sub>2</sub> concentration distribution and vertical velocity at Nanjing site show  
671 differences. The O<sub>2</sub> concentrations below boundary layer is substantially lower than that of  
672 Shanghai and Hangzhou. Besides, from August 7 to 10, vertical velocity is found extremely weak  
673 near the ground, as massive zero velocity lines appear below 500 m. From August 11 to 12,  
674 although the O<sub>3</sub> concentration is increasing as the high center is upper at the altitude of around 1  
675 km. But, the vertical movement below 2 km is dominated by upward airflow, thus the surface O<sub>3</sub>  
676 pollution is relieved with the O<sub>3</sub> concentration lower than the national standard (Fig. 2).

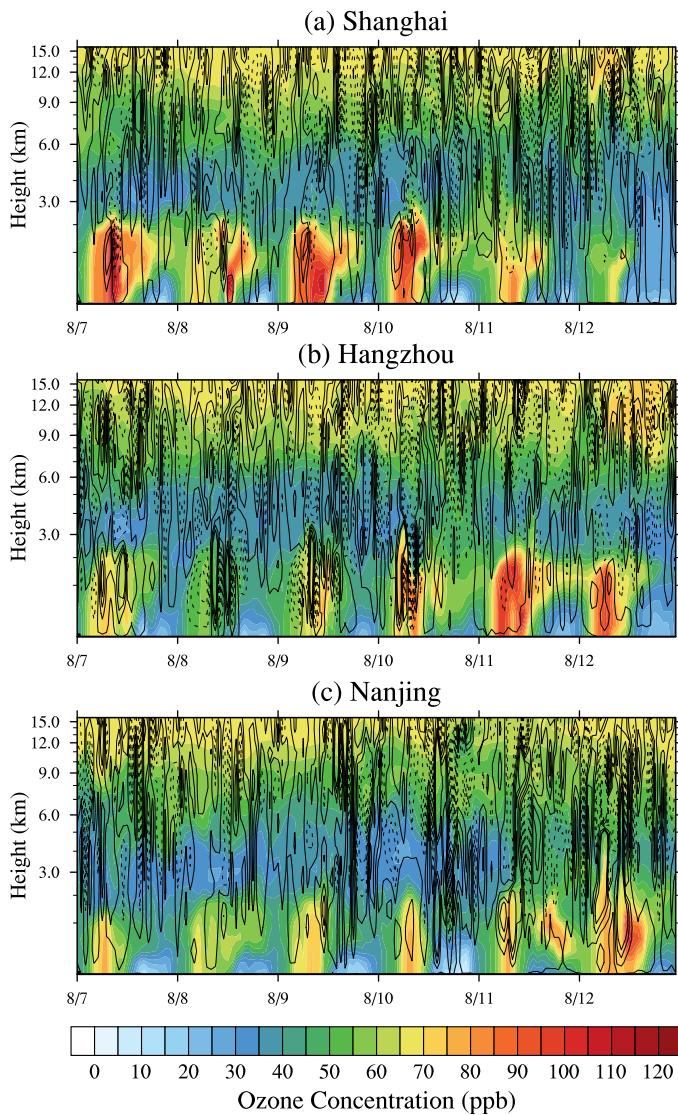
带格式的：下标

677

678



679



680

681 **Fig. 789.** Temporal variations of the vertical wind velocity and the vertical distribution of O<sub>3</sub> concentrations  
 682 above (a) Shanghai, (b) Hangzhou and (c) Nanjing during August 7 to 12 2013. The dotted lines show the  
 683 negative wind speeds and represent the downward airflows, while the solid lines show the zero-velocity lines  
 684 or the positive wind speeds— and zero vertical velocity represent upward airflow, while the dotted lines show  
 685 the negative wind speeds and represent downward air flow (an The interval of is 0.005 m/s).

686

687 It is also should be mentioned that the near-surface vertical velocities around these cities are  
 688 much lower than those at higher altitudes (Fig.8). Especially in the planetary boundary layer (< 1  
 689 km), lots of zero-velocity lines appear near the ground. This phenomenon may be related with the  
 690 upward airflow caused by Urban Heat Islands. Thus, the maximum centers of O<sub>3</sub> occur near the  
 691 surface below 500 m, and the vertical diffusion process plays a more important role in the

带格式的: 下标



带格式的: 下标

带格式的: 下标

692 | [accumulation of surface O<sub>3</sub>. The essential role of the vertical diffusion process in the O<sub>3</sub> episode is](#)  
693 | [similar to that reported by Zhu et al. \(2015\).](#)

694 | [Shanghai \(SH\) during August 7 to 12, 2013.](#)

695 |

#### 696 | 4.3 Process analysis for ozone formation

##### 697 | 4.3.1 Typical cities in the YRD region

698 | Fig. 8109 shows the ~~daily daytime~~ contributions of different atmospheric processes to the  
699 | formation of O<sub>3</sub> in Shanghai (SH), Nanjing (NJ), and Hangzhou (HZ) at the first modeling layer  
700 | from [August 4](#) to 15 ~~August~~-2013. As shown in the figure, for all cities during this period, the  
701 | major contributors to high O<sub>3</sub> concentrations include the vertical diffusion (VDIF), the dry  
702 | deposition (DDEP), the gas-phase chemistry (CHEM), and the total advection (TADV). TADV is  
703 | the sum of the horizontal advection (HADV) and the vertical advection (ZADV). In this study,  
704 | HADV and ZADV are considered together as TADV because they are inevitably linked as the  
705 | inseparable parts of air circulation. As discussed in Sect. 3.2, the strong sinking air causes slow  
706 | wind on the ground and little clouds in the sky, so the contributions of horizontal diffusion (HDIF)  
707 | and cloud processes (CLDS) are quite small during this episode.

708 | In the first layer of the urban areas of Shanghai (Fig. [1098a](#)), the averaged ~~daily daytime~~  
709 | contributions ~~from during August 4-15 August for~~ the vertical diffusion (VDIF), the gas-phase  
710 | chemistry (CHEM), the advection processes (TADV), and the dry deposition (DDEP) ~~during the~~  
711 | [daytime of August 4-15](#) are 9.95, 10.10, -11.74 and -7.28 ppb/h, respectively. Obviously, VDIF  
712 | and CHEM exhibit significant positive contributions to O<sub>3</sub> during most days, while TADV and  
713 | DDEP mainly show the consumption contributions. The sinking air caused by the weather system  
714 | discussed in Sect. 3.2 can trap heat and air pollutants on the ground, and ~~results in~~ VDIF ~~to~~  
715 | be the most import source of surface O<sub>3</sub>. Meanwhile, the hotter and dryer weather with more  
716 | sunshine, ~~above 40 °C and comparatively low relative humidity (shown in Fig. 4)~~, which is related  
717 | with the sinking air, can enhance the photochemical reactions. So, CHEM can form more O<sub>3</sub> on  
718 | the ground. Compared with the time series of CHEM and DDEP in which there are no obvious  
719 | fluctuations, the values of VDIF and TADV significantly change with the time, with the ~~daily~~  
720 | ~~daytime~~ mean contributions varying from 3.99 to 28.45 ppb/h for VDIF and from -2.56 to -28.13  
721 | ppb/h for TADV. These time variations should be related with the changes of vertical air

722 movement. For example, the value of VDIF on August 8 is only 3.99 ppb/h, which can be  
723 attributed to the local transient upward airflow over Shanghai (shown in Fig. 98a7). On August 10,  
724 however, VDIF can contribute 28.45 ppb O<sub>3</sub> per hour, which may be related with the enhanced  
725 downward air movement caused by the peripheral circulation of Typhoon Utor. Moreover, during  
726 the high-level O<sub>3</sub> episode from August 7-12, the mean values for VDIF, CHEM, TADV and DDEP  
727 are 13.41, 11.21, -8.37 and -14.74 ppb/h. But after August 12, the mean contributions of VDIF,  
728 CHEM, TADV and DDEP decrease to 5.35, 9.53, -5.52 and -10.85 ppb/h. These reductions should  
729 be related with the process that the subtropical high moves eastward and northward forced by  
730 Typhoon Utor (Fig. 5d). By quantifying the relative importance of each process to O<sub>3</sub> formation,  
731 the IPR analysis provides a fundamental explanation for the synthetically influence of the high  
732 pressure and the typhoon system, which has been discussed in Sect. 3.2 and 4.1, and further  
733 illustrates the exact mechanism.

734 Fig. 8109b presents the result of IPR analysis for Hangzhou. During August 4-15, VDIF and  
735 CHEM are the major source of surface O<sub>3</sub> with the average contribution of 5.36 ppb/h for VDIF  
736 and 10.97 ppb/h for CHEM, while TADV and DDEP are two important sinks for O<sub>3</sub> with the  
737 average contribution of -9.63 ppb/h for TADV and -5.14 ppb/h for DDEP. Synthetically impacted  
738 by Western Pacific subtropical high and Typhoon Utor, the mean contributions during the O<sub>3</sub>  
739 episode (from August 7 to August 12) for VDIF, CHEM, TADV and DDEP increase to 7.21, 12.61,  
740 -11.51 and -5.92 ppb/h, respectively. The highest VDIF contribution occurs on August 10-11, and  
741 the over-standard of O<sub>3</sub> concentration appears on August 10-12 as well, which may be attributed to  
742 the effect of typhoon's peripheral circulation, implying Typhoon Utor also plays an essential role  
743 in the formation of O<sub>3</sub> pollution in Hangzhou. After Typhoon Utor approaches close enough to  
744 Hangzhou, the wind direction is mainly dominated by the southeast wind (Fig. 4b), and the mean  
745 values of VDIF, CHEM, TADV and DDEP finally decrease to 4.84, 10.08, -8.92 and -4.78 ppb/h,  
746 respectively. In a word, Hangzhou is located close to Shanghai, so the temporal variations of VDIF,  
747 CHEM, TADV and DDEP in Hangzhou are similar to those in Shanghai.

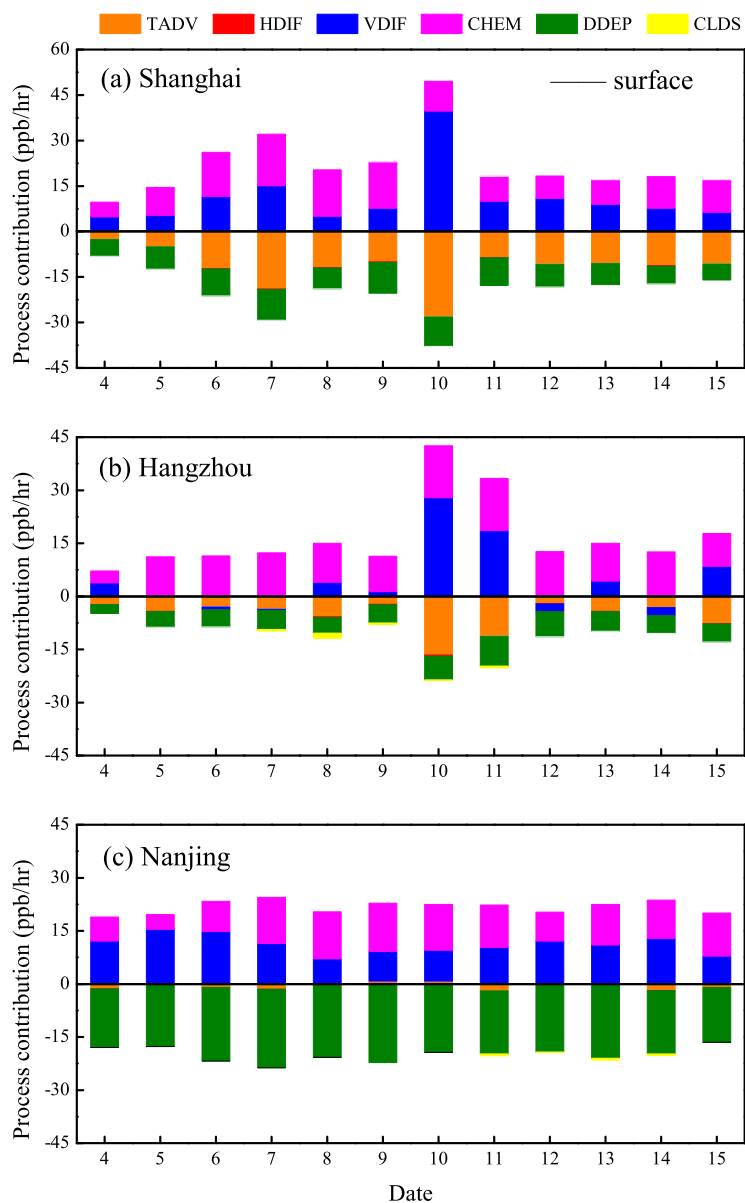
748 However, the similar variation pattern of VDIF, CHEM, TADV and DDEP occurring in  
749 Shanghai and Hangzhou does not appear in Nanjing. As shown in Fig. 8109c, the mean  
750 contributions of VDIF, CHEM, TADV and DDEP to surface O<sub>3</sub> in Nanjing are 11.31, 9.55 -1.34  
751 and -17.57 ppb/h during the whole period, while the values during 7-12 August are 10.32, 10.70,



752 -0.99 and -18.42 ppb/h. There are no apparent fluctuations or sudden increases of these  
753 contributors during the period from August 4 to 15, so are the O<sub>3</sub> concentration (Fig. 2),  
754 temperature and relative humidity (Fig. 4a). implying Nanjing is generally under the control of the  
755 Western Pacific subtropical high and can hardly be affected by the typhoon system. The above  
756 discussion in Section 4.2 also proved this conclusion. As a typical city in the northwest inland area  
757 of the YRD region (NIR), Nanjing is located far away from the sea, which means it may not be  
758 easily affected by the peripheral circulation of the typhoon system~~weather system from the ocean.~~

759 Additionally, at the altitude of 500 m and 1500 m above Shanghai, Nanjing, and Hangzhou  
760 (not shown), CHEM is also the major contributor to O<sub>3</sub> formation, with the values a litter lower  
761 than those at the surface, suggesting that there are strong photochemical reactions in the whole  
762 boundary layer of these YRD cities. In contrast, VDIF has an opposite effect in the middle of the  
763 boundary layer, with the negative contributions for O<sub>3</sub> of -3.26 ppb/h in Shanghai, -2.37 ppb/h in  
764 Hangzhou, and -3.21 ppb/h in Nanjing, respectively (not shown). The loss of O<sub>3</sub> at higher  
765 atmospheric level caused by VDIF further proves the essential role of the downward vertical  
766 movement in this O<sub>3</sub> episode.

767



768

769 **Fig. 8910.** Variations of the ~~daily daytime~~ mean values for the contributions of individual processes to O<sub>3</sub>  
 770 formation in (a) Shanghai, (b) Hangzhou, and (c) Nanjing from ~~August 4 to 15 August~~ 2013 at the surface  
 771 layer. The contributors include the total advection (TADV), the horizontal diffusion (HDIF), the vertical  
 772 diffusion (VDIF), the gas-phase chemistry (CHEM), the dry deposition (DDEP), and the cloud processes  
 773 with the aqueous chemistry (CLDS).

774

775 **4.3.2 Spatial distribution of the contributors for the O<sub>3</sub> episode over the YRD region**

776 Fig. 9-10 demonstrates the spatial distribution of the mean contributions of main processes

777 (VDIF, CHEM, DDEP and TADV) to the [ozone formation of this high-level O<sub>3</sub> episode](#) at the  
778 lowest modeling layer in domain 3 [during this high-level O<sub>3</sub> episode](#). The modeling results from 7  
779 to 12 August are averaged to provide the mean values.

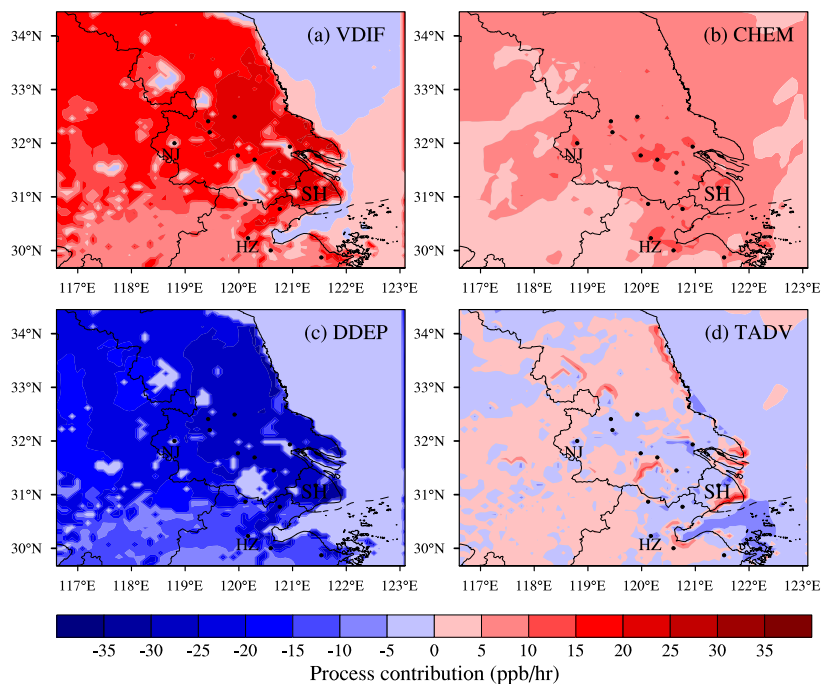
780 Similar to the results shown in Fig. [89](#), Fig. [910](#) illustrates that the vertical diffusion (VDIF)  
781 and the gas-phase chemistry (CHEM) exhibit significant positive contributions to O<sub>3</sub> over the  
782 YRD region and the surrounding areas during the high-level O<sub>3</sub> episode. The contributions of  
783 VDIF in domain 3 (Fig. [109a](#)) range from 5 to 25 ppb/h, with the high values (> 20 ppb/h)  
784 occurring in the southeast coastal areas. For CHEM (Fig. [109b](#)), the contributions vary within the  
785 range of 0-15 ppb/h, with the high values over 10 ppb/h appearing in and around the big cities. As  
786 discussed above, these regional positive contributions of VDIF and CHEM over domain 3 should  
787 be related to the facts that the whole region is under the control of the Western Pacific subtropical  
788 high. With respect to the higher contributions of CHEM in the urban areas, they should be  
789 attributed to the spatial distribution of the emissions of O<sub>3</sub> precursors, which is also higher in the  
790 cities. Furthermore, higher air temperature in the cities related with the urban heat island may  
791 enhance the chemical reactions and form more O<sub>3</sub> in these areas as well.           

792  
793 [As shown in Fig. 9e, For DDEP, it](#) is the main critical factor of the consumption of O<sub>3</sub>, with  
794 the negative contributions varying from 0 to -25 ppb/h over the modeling domain 3 [\(Fig. 10c\)](#).  
795 Small values usually occur on the water, which may be related with less air pollution over rivers,  
796 lakes and oceans. High values can be found on land, especially in the southeast coastal areas. For  
797 the contributions of TADV [\(Fig. 9d\)](#), the values in domain 3 range from -10 to 10 ppb/h, with the  
798 positive contributions generally occurring on land while the negative [\(consuming\)](#)-ones appearing  
799 on the water [\(Fig. 10d\)](#). The maximum positive contributions of TADV are usually found along  
800 the boundary between the land and the water, which should be explained by the facts that the  
801 land-sea breeze circulations can play an important role in the redistribution of the formed O<sub>3</sub>. [On](#)  
802 [account of the high-pressure system and so-caused sinking airflows in the YRD region, the](#)  
803 [background wind is relatively weak in comparison to the local atmospheric circulation, thus the](#)  
804 [sea breeze can easily bring more generated O<sub>3</sub> to the seashore.](#)

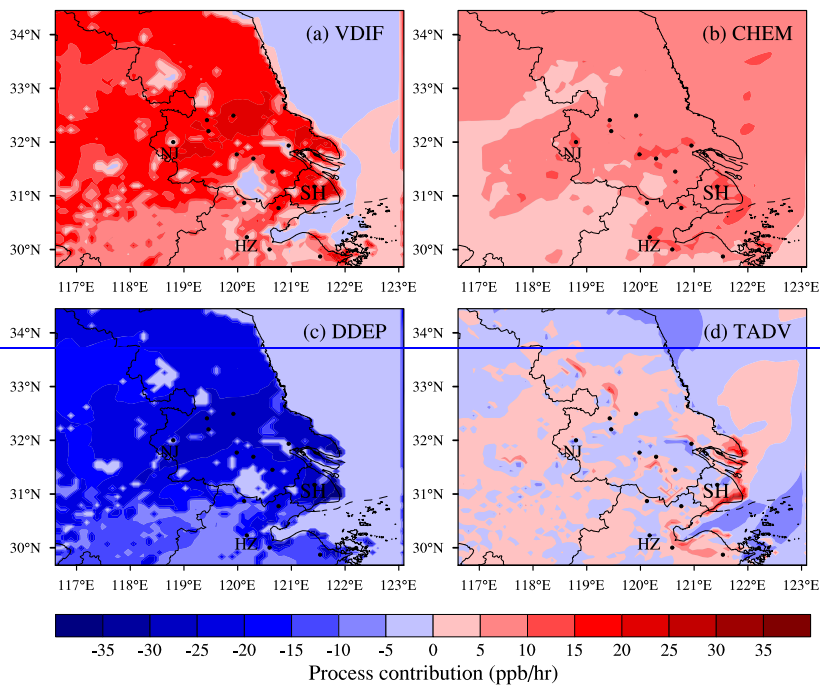
805 [From the discussion in Section 3 and 4.2](#)

806 [In all, more active photochemical reactions and the vertical diffusion play a significant role in](#)

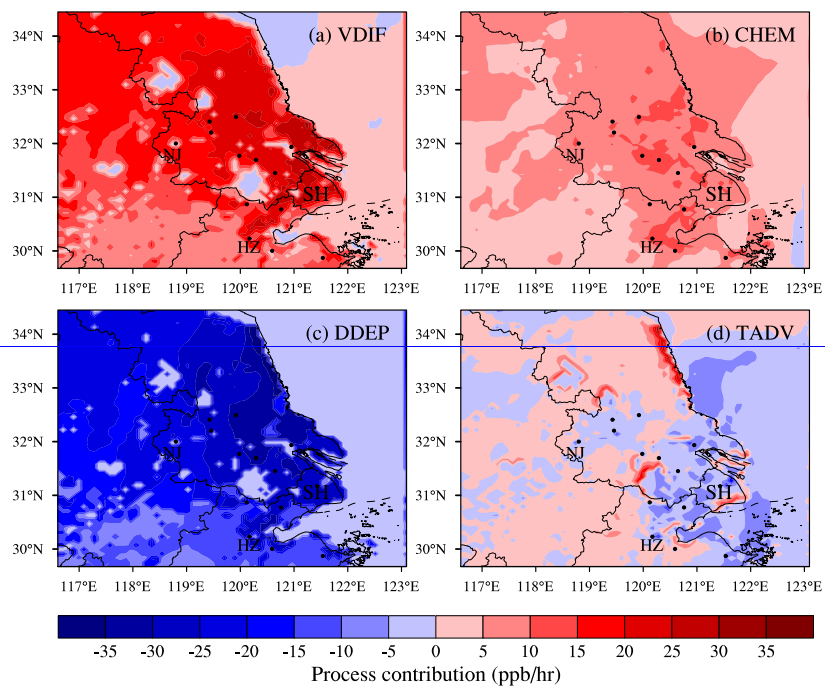
807 ~~the accumulation of surface  $O_3$ , and lead to the high-level  $O_3$  pollution episode over the YRD~~  
 808 ~~region. The major driving factor should be the Western Pacific subtropical high. Moreover, the~~  
 809 ~~changes in the contributions of VDIF, CHEM, DDEP, and TADV between August 7-9 and August~~  
 810 ~~10-12 CHEM exhibit a similar spatial pattern with the high values and increased values mostly~~  
 811 ~~concentrate concentrating in the southeast coastal areas (Fig. 12), implying the Typhoon Utor also~~  
 812 ~~plays a collaborative effect, which is clearly shown in Fig. 13. The details and the processes of the~~  
 813 ~~synthetical effects of high pressure and typhoon system have been discussed in Sect. 3.2, 4.2 and~~  
 814 ~~4.3.1 as well.~~  
 815  
 816



817  
 818 **Fig. 9. The contributions of main processes to  $O_3$  formation over the YRD region, including (a) vertical**  
 819 **diffusion (VDIF), (b) gas chemistry (CHEM), (c) dry deposition (DDEP), and (d) total advection (TADV).**  
 820 **The values are averaged from August 7 to 12 2013.**  
 821



822



823

824 [Fig. 12. The daytime contributions of main processes to  \$O\_3\$  formation over the YRD region, including \(a\)](#)  
 825 [vertical diffusion \(VDIF\), \(b\) gas chemistry \(CHEM\), \(c\) dry deposition \(DDEP\), and \(d\) total advection](#)

826 | (TADV). The values are averaged from August 10 to 12 2013.

827

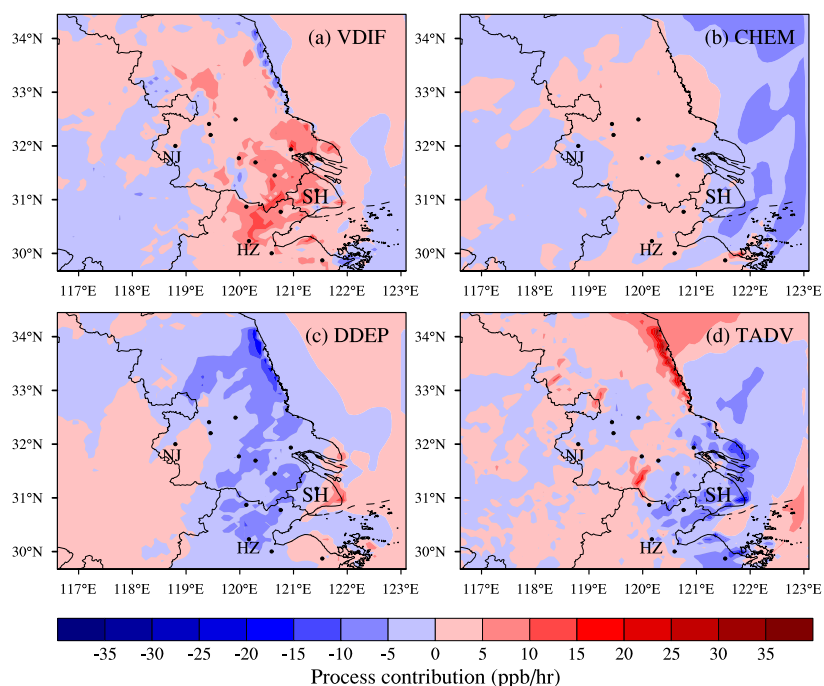
828 | From the discussion in Section 3 and 4.2, it can be deduced that typhoon Utor plays an  
829 | important role in the formation of ozone over the YRD region during August 10-12. To clearly  
830 | clarify the effect of the typhoon system in this O<sub>3</sub> pollution episode, we firstly average the  
831 | modeling results of VDIF, CHEM, DDEP and TADV during August 10-12 to show their  
832 | contributions to O<sub>3</sub> formation when the typhoon system plays an important role. Secondly, the  
833 | modeling results of these processes from August 7 to 9 are also averaged to provide their  
834 | contributions when only the subtropical high dominates the episode. Finally, the differences of the  
835 | contributions of VDIF, CHEM, DDEP and TADV between the period of August 7-9 and August  
836 | 10-12 are calculated to reveal the role of the typhoon system in this severe high O<sub>3</sub> episode (Fig.  
837 | 11). As shown in Fig. 11a, when YRD is affected by the peripheral circulation of Typhoon Utor,  
838 | the contributions of VDIF over the YRD region increase by 0-15 ppb/h, with the higher increment  
839 | values (> 30 ppb/h) occurring in the southeast coastal region (SCR) and center inland region (CIR),  
840 | implying that SCR and CIR can be largely affected by the peripheral subsidence airflows of the  
841 | typhoon system. As to the contributions of CHEM, the increases caused by the typhoon system are  
842 | 0-5 ppb/h over the YRD region, and the higher increment also appears in the coastal areas (Fig.  
843 | 11b). For DDEP, influenced by typhoon Utor, its negative contributions decrease up to -20 ppb/h,  
844 | with the largest reduction along the coastline (Fig. 11c). For TADV, with the approaching of  
845 | typhoon Utor, the contributions of TADV particularly decrease by 0-20 ppb/h, especially in the  
846 | southeast coastal region (Fig. 11d).

847 | In all, during this high-level O<sub>3</sub> pollution episode, more active photochemical reactions and  
848 | the vertical diffusion play a significant role in the accumulation of surface O<sub>3</sub> over the YRD region.  
849 | The major driving factor should be the Western Pacific subtropical high. Moreover, the changes in  
850 | the contributions of VDIF, CHEM, DDEP, and TADV between August 7-9 and August 10-12  
851 | exhibit a similar spatial pattern with the high values mostly concentrating in the southeast coastal  
852 | areas (Fig. 12), implying the Typhoon Utor also plays a collaborative effect.

853

带格式的：下标

带格式的：下标



854  
 855 **Fig. 131.** The difference of daytime mean contributions of main processes to O<sub>3</sub> formation over the YRD  
 856 region between the period of August 10-12 and August 7-9, including (a) vertical diffusion (VDIF),  
 857 (b) gas chemistry (CHEM), (c) dry deposition (DDEP), and (d) total advection (TADV) between the period  
 858 of August 7-9 and August 10-12 (values averaged from August 10-12 minus that of August 7-9).

859  
 860 **5. Conclusions**

861 ~~In this study, By means of observational analysis and numerical simulation,~~ the  
 862 characteristics and the essential impact factors of a typical regional continuous O<sub>3</sub> pollution over  
 863 the YRD region is investigated ~~by means of observational analysis and numerical simulation. Base~~  
 864 ~~on the observation data, it is found that this high level O<sub>3</sub> The~~ episode lasted ~~eds~~ for nearly a week  
 865 from ~~August 7 to 12 August~~ 2013, with the O<sub>3</sub> concentration exceeding the national air quality  
 866 standard in more than half of the cities over the YRD region. ~~In the cities of Jiaxing, Changzhou~~  
 867 ~~and Nantong, high O<sub>3</sub> concentrations can reach the values over 160 ppb. Fine weather conditions,~~  
 868 ~~such as extremely high temperature, low relative humidity, and weak wind speed, provide a~~  
 869 ~~favorable atmospheric environment for the complicated photochemical reactions and help to form~~  
 870 O<sub>3</sub>. The analysis of weather systems and the modeling results from WRF/CMAQ all illustrate that  
 871 the continuous strong Western Pacific subtropical high is the leading factor of the abnormally high

872 temperature weather and the heavy O<sub>3</sub> pollution, by inducing more sinking air to trap heat as well  
873 as air pollutants at the surface. ~~Meanwhile, the~~ development of this episode is closely related to  
874 the movement of Typhoon Utor as well. The temporal variations of the vertical wind velocity and  
875 O<sub>3</sub> concentrations show that when the YRD region is at the front of moving typhoon system, the  
876 downward airflow is enhanced in the boundary layer with fine weather, and thereby the air  
877 pollutants are trapped and accumulated near the surface. Moreover, in the last stage of the O<sub>3</sub>  
878 episode, the activity of Typhoon Utor weakens the strength of the subtropical high and forces it to  
879 retreat easterly and move northward, and the prevailing southeasterly surface wind related with the  
880 approaching of Typhoon Utor contributes to the mitigation of the O<sub>3</sub> pollution.

881 The Integrated Process Rate (IPR) analysis implemented in CMAQ is specially carried out to  
882 quantify the relative contributions of individual processes and give a fundamental explanation.  
883 ~~Over the YRD region, during the high-level O<sub>3</sub> episode from August 7-12, the vertical diffusion~~  
884 (VDIF) and the gas-phase chemistry (CHEM) exhibit significant positive contributions to surface  
885 O<sub>3</sub>; ~~over the YRD region, with the high values over 20 ppb/h for VDIF and over 10 ppb/h for~~  
886 CHEM. ~~The dry deposition (DDEP) is the major sink of surface O<sub>3</sub>, while the total advection~~  
887 (TADV) can give the positive contribution on land and the negative contribution on the water. ~~The~~  
888 ~~dry deposition (DDEP) is the major sink of surface O<sub>3</sub>, while the contributions of horizontal~~  
889 ~~diffusion (HDIF) and cloud processes (CLDS) are quite small. To some extent, the distribution~~  
890 ~~pattern reflects the heterogeneity of emissions and the effects of weather system. Influenced by the~~  
891 ~~sinking air as well as the fine weather induced by the Western Pacific subtropical high, the~~  
892 ~~contributions of VDIF and CHEM to surface O<sub>3</sub> maintain the high values of 13.41 and 11.21 ppb/h~~  
893 ~~for Shanghai, 7.21 and 12.61 ppb/h for Hangzhou, and 10.32 and 10.70 ppb/h for Nanjing,~~  
894 ~~respectively. Moreover, on August 10-12, the YRD region cities close to the sea are is~~ apparently  
895 affected by the periphery circulation of Typhoon Utor, with the contributions of VDIF over the  
896 YRD region increasing by 0-15 ppb/h, the contributions of CHEM increasing by 0-5 ppb/h, and  
897 the contributions of DDEP and TADV decreasing. Especially in the coastal cities, such as  
898 Shanghai and Hangzhou, the effects of the typhoon system are more obvious, particularly an  
899 increase to 28.45 ppb/h in Shanghai and 19.76 ppb/h in Hangzhou. Meanwhile, the contributions  
900 of CHEM in YRD increase by 0-5 ppb/h, while the negative contributions of DDEP decrease by  
901 0-20 ppb/h with large decrease values appear along the coastline. In contrast, the cities in the



902 ~~northwest inland area~~ of the YRD region, ~~(NIR)~~ which are far away from the sea, ~~are generally~~  
903 ~~under the control of the subtropical high and can hardly be affected by the typhoon system owing~~  
904 ~~to the location far away from the sea. In the end, W~~when the typhoon system significantly  
905 weakens the high pressure system, the contributions of VDIF, CHEM, TADV, and DDEP decrease  
906 to a low level in all cities.

907 WRF-CMAQ model system shows a relatively good performance in simulation of the O<sub>3</sub>  
908 episode, with the simulated meteorological conditions and air pollutant concentrations basically in  
909 agreement with the observations in most YRD cities. Our results in this study can provide an  
910 insight for the formation mechanism of regional O<sub>3</sub> pollution in East Asia, and help to forecast the  
911 O<sub>3</sub> pollution synthetically impacted by the Western Pacific subtropical high and the tropical  
912 cyclone system.

913

#### 914 **Acknowledgments**

915 This study was supported by the National Natural Science Foundation of China (41475122,  
916 91544230, 41575145), the National Special Fund for Environmental Protection Research in the  
917 Public Interest (201409008), EU 7th Framework Marie Curie Actions IRSES project REQUA  
918 (PIRSES-GA-2013-612671), and the National Science Foundation of Jiangsu Province  
919 (BE2015151). The authors would like to thank [Hongli Wang from Shanghai Academy of](#)  
920 [Environmental Sciences for providing VOCs observation data](#), Xiaoxun Xie for preliminary data  
921 processing, and the anonymous reviewers for their constructive and precious comments on this  
922 manuscript.

923

#### 924 **References**

- 925 An, X., Zhu, T., Wang, Z., Li, C., and Wang, Y.: A modeling analysis of a heavy air pollution episode occurred in  
926 Beijing, *Atmos Chem Phys*, 7, 3103-3114, 2007.
- 927 Byun, D., and Schere, K. L.: Review of the governing equations, computational algorithms, and other components  
928 of the models-3 Community Multiscale Air Quality (CMAQ) modeling system, *Appl Mech Rev*, 59, 51-77,  
929 10.1115/1.2128636, 2006.
- 930 Chan, C. K., and Yao, X.: Air pollution in mega cities in China, *Atmos Environ*, 42, 1-42,  
931 10.1016/j.atmosenv.2007.09.003, 2008.
- 932 Chen, F., and Dudhia, J.: Coupling an advanced land surface-hydrology model with the Penn State-NCAR MM5  
933 modeling system. Part I: Model implementation and sensitivity, *Mon Weather Rev*, 129, 569-585, Doi  
934 10.1175/1520-0493(2001)129<0569:Caalsh>2.0.Co;2, 2001.

935 Chen, D., Zhou, B., Beirle, S., Chen, L. M., and Wagner, T.: Tropospheric NO<sub>2</sub> column densities deduced from  
936 zenith-sky DOAS measurements in Shanghai, China, and their application to satellite validation, *Atmos Chem*  
937 *Phys*, 9, 3641-3662, 2009.

938 Cheng, W. L., Lai, L. W., Den, W., Wu, M. T., Hsueh, C. A., Lin, P. L., Pai, C. L., and Yan, Y. L.: The  
939 relationship between typhoons' peripheral circulation and ground-level ozone concentrations in central Taiwan,  
940 *Environ Monit Assess*, 186, 791-804, 10.1007/s10661-013-3417-7, 2014.

941 Chiqueto, J., Silva, M. E. S.: São Paulo "Surface Ozone Layer" and the atmosphere: characteristics of tropospheric  
942 ozone concentrations in the city and how the atmosphere influences them, VDM Verlag Dr. Muller,  
943 Saarbrücken, 2010.

944 Ding, A. J., Fu, C. B., Yang, X. Q., Sun, J. N., Zheng, L. F., Xie, Y. N., Herrmann, E., Nie, W., Petaja, T.,  
945 Kerminen, V. M., and Kulmala, M.: Ozone and fine particle in the western Yangtze River Delta: an overview  
946 of 1 yr data at the SORPES station, *Atmos Chem Phys*, 13, 5813-5830, 10.5194/acp-13-5813-2013, 2013.

947 Duan, J. C., Tan, J. H., Yang, L., Wu, S., and Hao, J. M.: Concentration, sources and ozone formation potential of  
948 volatile organic compounds (VOCs) during ozone episode in Beijing, *Atmos Res*, 88, 25-35,  
949 10.1016/j.atmosres.2007.09.004, 2008.

950 Emmons, L. K., Walters, S., Hess, P. G., Lamarque, J. F., Pfister, G. G., Fillmore, D., Granier, C., Guenther, A.,  
951 Kinnison, D., Laepple, T., Orlando, J., Tie, X., Tyndall, G., Wiedinmyer, C., Baughcum, S. L., and Kloster, S.:  
952 Description and evaluation of the Model for Ozone and Related chemical Tracers, version 4 (MOZART-4),  
953 *Geosci Model Dev*, 3, 43-67, 2010.

954 Fann, N., and Risley, D.: The public health context for PM<sub>2.5</sub> and ozone air quality trends, *Air Qual Atmos Hlth*, 6,  
955 1-11, 10.1007/s11869-010-0125-0, 2013.

956 Feng, Z. W., Jin, M. H., Zhang, F. Z., and Huang, Y. Z.: Effects of ground-level ozone (O<sub>3</sub>) pollution on the  
957 yields of rice and winter wheat in the Yangtze River Delta, *J Environ Sci-China*, 15, 360-362, 2003.

958 Foley, K. M., Roselle, S. J., Appel, K. W., Bhave, P. V., Pleim, J. E., Otte, T. L., Mathur, R., Sarwar, G., Young, J.  
959 O., Gilliam, R. C., Nolte, C. G., Kelly, J. T., Gilliland, A. B., and Bash, J. O.: Incremental testing of the  
960 Community Multiscale Air Quality (CMAQ) modeling system version 4.7, *Geosci Model Dev*, 3, 205-226,  
961 2010.

962 Gao, J. H., Bin, Z., Xiao, H., Kang, H. Q., Hou, X. W., and Shao, P.: A case study of surface ozone source  
963 apportionment during a high concentration episode, under frequent shifting wind conditions over the Yangtze  
964 River Delta, China, *Sci Total Environ*, 544, 853-863, 10.1016/j.scitotenv.2015.12.039, 2016.

965 Geng, F. H., Tie, X. X., Xu, J. M., Zhou, G. Q., Peng, L., Gao, W., Tang, X., and Zhao, C. S.: Characterizations of  
966 ozone, NO<sub>x</sub>, and VOCs measured in Shanghai, China, *Atmos Environ*, 42, 6873-6883,  
967 10.1016/j.atmosenv.2008.05.045, 2008.

968 Goncalves, M., Jimenez-Guerrero, P., and Baldasano, J. M.: Contribution of atmospheric processes affecting the  
969 dynamics of air pollution in South-Western Europe during a typical summertime photochemical episode,  
970 *Atmos Chem Phys*, 9, 849-864, 2009.

971 Grell, G. A., and Devenyi, D.: A generalized approach to parameterizing convection combining ensemble and data  
972 assimilation techniques, *Geophys Res Lett*, 29, ArtId 169310, 10.1029/2002gl015311, 2002.

973 Guo, H., Jiang, F., Cheng, H. R., Simpson, I. J., Wang, X. M., Ding, A. J., Wang, T. J., Saunders, S. M., Wang, T.,  
974 Lam, S. H. M., Blake, D. R., Zhang, Y. L., and Xie, M.: Concurrent observations of air pollutants at two sites  
975 in the Pearl River Delta and the implication of regional transport, *Atmos Chem Phys*, 9, 7343-7360, 2009.

976 Han, S. Q., Bian, H., Feng, Y. C., Liu, A. X., Li, X. J., Zeng, F., and Zhang, X. L.: Analysis of the Relationship  
977 between O<sub>3</sub>, NO and NO<sub>2</sub> in Tianjin, China, *Aerosol Air Qual Res*, 11, 128-139, 10.4209/aaqr.2010.07.0055,  
978 2011.

979 Hong, S. Y., Dudhia, J., and Chen, S. H.: A revised approach to ice microphysical processes for the bulk  
980 parameterization of clouds and precipitation, *Mon Weather Rev*, 132, 103-120, Doi  
981 10.1175/1520-0493(2004)132<0103:Aratim>2.0.Co;2, 2004.

982 Hong, S. Y., Noh, Y., and Dudhia, J.: A new vertical diffusion package with an explicit treatment of entrainment  
983 processes, *Mon Weather Rev*, 134, 2318-2341, Doi 10.1175/Mwr3199.1, 2006.

984 Huang, J. P., Fung, J. C. H., Lau, A. K. H., and Qin, Y.: Numerical simulation and process analysis of  
985 typhoon-related ozone episodes in Hong Kong, *J Geophys Res-Atmos*, 110, Artn  
986 D0530110.1029/2004jd004914, 2005.

987 Huang, J. P., Fung, J. C. H., and Lau, A. K. H.: Integrated processes analysis and systematic meteorological  
988 classification of ozone episodes in Hong Kong, *J Geophys Res-Atmos*, 111, Artn  
989 D2030910.1029/2005jd007012, 2006.

990 Hung, C. H., and Lo, K. C.: Relationships between Ambient Ozone Concentration Changes in Southwestern  
991 Taiwan and Invasion Tracks of Tropical Typhoons, *Adv Meteorol, Artn 40297610.1155/2015/402976*, 2015.

992 Jenkin, M. E., and Clemitshaw, K. C.: Ozone and other secondary photochemical pollutants: chemical processes  
993 governing their formation in the planetary boundary layer, *Atmos Environ*, 34, 2499-2527, Doi  
994 10.1016/S1352-2310(99)00478-1, 2000.

995 Jiang, F., Wang, T. J., Wang, T. T., Xie, M., and Zhao, H.: Numerical modeling of a continuous photochemical  
996 pollution episode in Hong Kong using WRF-chem, *Atmos Environ*, 42, 8717-8727,  
997 10.1016/j.atmosenv.2008.08.034, 2008.

998 Jiang, F., Zhou, P., Liu, Q., Wang, T. J., Zhuang, B. L., and Wang, X. Y.: Modeling tropospheric ozone formation  
999 over East China in springtime, *J Atmos Chem*, 69, 303-319, 10.1007/s10874-012-9244-3, 2012.

1000 Kim, H. J., and Wang, B.: Sensitivity of the WRF Model Simulation of the East Asian Summer Monsoon in 1993  
1001 to Shortwave Radiation Schemes and Ozone Absorption, *Asia-Pac J Atmos Sci*, 47, 167-180,  
1002 10.1007/s13143-011-0006-y, 2011.

1003 Lam, K. S., Wang, T. J., Wu, C. L., and Li, Y. S.: Study on an ozone episode in hot season in Hong Kong and  
1004 transboundary air pollution over Pearl River Delta region of China, *Atmos Environ*, 39, 1967-1977,  
1005 10.1016/j.atmosenv.2004.11.023, 2005.

1006 Landry, J. S., Neilson, E. T., Kurz, W. A., and Percy, K. E.: The impact of tropospheric ozone on landscape-level  
1007 merchantable biomass and ecosystem carbon in Canadian forests, *Eur J Forest Res*, 132, 71-81,  
1008 10.1007/s10342-012-0656-z, 2013.

1009 Li, L., Chen, C. H., Huang, C., Huang, H. Y., Zhang, G. F., Wang, Y. J., Chen, M. H., Wang, H. L., Chen, Y. R.,  
1010 Streets, D. G., and Fu, J. M.: Ozone sensitivity analysis with the MM5-CMAQ modeling system for Shanghai,  
1011 *J Environ Sci-China*, 23, 1150-1157, 10.1016/S1001-0742(10)60527-X, 2011.

1012 Li, L., Chen, C. H., Huang, C., Huang, H. Y., Zhang, G. F., Wang, Y. J., Wang, H. L., Lou, S. R., Qiao, L. P.,  
1013 Zhou, M., Chen, M. H., Chen, Y. R., Streets, D. G., Fu, J. S., and Jang, C. J.: Process analysis of regional  
1014 ozone formation over the Yangtze River Delta, China using the Community Multi-scale Air Quality modeling  
1015 system, *Atmos Chem Phys*, 12, 10971-10987, 10.5194/acp-12-10971-2012, 2012.

1016 Li, M. M., Song, Y., Huang, X., Li, J. F., Mao, Y., Zhu, T., Cai, X. H., and Liu, B.: Improving mesoscale modeling  
1017 using satellite-derived land surface parameters in the Pearl River Delta region, China, *J Geophys Res-Atmos*,  
1018 119, 6325-6346, 10.1002/2014JD021871, 2014.

1019 Li, M. M., Song, Y., Mao, Z. C., Liu, M. X., and Huang, X.: Impacts of thermal circulations induced by  
1020 urbanization on ozone formation in the Pearl River Delta region, China, *Atmos Environ*, 127, 382-392,  
1021 10.1016/j.atmosenv.2015.10.075, 2016.

1022 Liao, J. B., Wang, T. J., Wang, X. M., Xie, M., Jiang, Z. Q., Huang, X. X., and Zhu, J. L.: Impacts of different

1023 urban canopy schemes in WRF/Chem on regional climate and air quality in Yangtze River Delta, China,  
1024 Atmos Res, 145, 226-243, 10.1016/j.atmosres.2014.04.005, 2014.

1025 Liao, J. B., Wang, T. J., Jiang, Z. Q., Zhuang, B. L., Xie, M., Yin, C. Q., Wang, X. M., Zhu, J. L., Fu, Y., and  
1026 Zhang, Y.: WRF/Chem modeling of the impacts of urban expansion on regional climate and air pollutants in  
1027 Yangtze River Delta, China, Atmos Environ, 106, 204-214, 10.1016/j.atmosenv.2015.01.059, 2015.

1028 Liu, X. H., Zhang, Y., Xing, J., Zhang, Q. A., Wang, K., Streets, D. G., Jang, C., Wang, W. X., and Hao, J. M.:  
1029 Understanding of regional air pollution over China using CMAQ, part II. Process analysis and sensitivity of  
1030 ozone and particulate matter to precursor emissions, Atmos Environ, 44, 3719-3727,  
1031 10.1016/j.atmosenv.2010.03.036, 2010.

1032 Liu, Q., Lam, K. S., Jiang, F., Wang, T. J., Xie, M., Zhuang, B. L., and Jiang, X. Y.: A numerical study of the  
1033 impact of climate and emission changes on surface ozone over South China in autumn time in 2000-2050,  
1034 Atmos Environ, 76, 227-237, 10.1016/j.atmosenv.2013.01.030, 2013.

1035 Lou, S. J., Liao, H., and Zhu, B.: Impacts of aerosols on surface-layer ozone concentrations in China through  
1036 heterogeneous reactions and changes in photolysis rates, Atmos Environ, 85, 123-138,  
1037 10.1016/j.atmosenv.2013.12.004, 2014.

1038 Lu, W. Z., and Wang, X. K.: Evolving trend and self-similarity of ozone pollution in central Hong Kong ambient  
1039 during 1984-2002, Sci Total Environ, 357, 160-168, 10.1016/j.scitotenv.2005.03.015, 2006.

1040 Ma, J. Z., Xu, X. B., Zhao, C. S., and Yan, P.: A review of atmospheric chemistry research in China:  
1041 Photochemical smog, haze pollution, and gas-aerosol interactions, Adv Atmos Sci, 29, 1006-1026,  
1042 10.1007/s00376-012-1188-7, 2012.

1043 Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., and Clough, S. A.: Radiative transfer for  
1044 inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave, J Geophys Res-Atmos,  
1045 102, 16663-16682, Doi 10.1029/97jd00237, 1997.

1046 Monin, A. S., Obukhov, A.M.: Basic laws of turbulent mixing in the surface layer of the atmosphere,  
1047 Contributions of the Geophysical Institute of the Slovak Academy of Sciences 151, 163-187, 1954.

1048 [Peng, J. B.: An Investigation of the Formation of the Heat Wave in Southern China in Summer 2013 and the](#)  
1049 [Relevant Abnormal Subtropical High Activities, Atmospheric & Oceanic Science Letters, 7, 286-290, 2014.](#)

1050 Ran, L., Zhao, C. S., Geng, F. H., Tie, X. X., Tang, X., Peng, L., Zhou, G. Q., Yu, Q., Xu, J. M., and Guenther, A.:  
1051 Ozone photochemical production in urban Shanghai, China: Analysis based on ground level observations, J  
1052 Geophys Res-Atmos, 114, Artn D1530110.1029/2008jd010752, 2009.

1053 Shao, M., Zhang, Y. H., Zeng, L. M., Tang, X. Y., Zhang, J., Zhong, L. J., and Wang, B. G.: Ground-level ozone  
1054 in the Pearl River Delta and the roles of VOC and NO<sub>x</sub> in its production, J Environ Manage, 90, 512-518,  
1055 10.1016/j.jenvman.2007.12.008, 2009.

1056 Shi, C. Z., Wang, S. S., Liu, R., Zhou, R., Li, D. H., Wang, W. X., Li, Z. Q., Cheng, T. T., and Zhou, B.: A study  
1057 of aerosol optical properties during ozone pollution episodes in 2013 over Shanghai, China, Atmos Res, 153,  
1058 235-249, 10.1016/j.atmosres.2014.09.002, 2015.

1059 Sillman, S.: The relation between ozone, NO<sub>x</sub> and hydrocarbons in urban and polluted rural environments, Atmos  
1060 Environ, 33, 1821-1845, Doi 10.1016/S1352-2310(98)00345-8, 1999.

1061 State Environmental Protection Administration of China: China National Environmental Protection Standard:  
1062 Automated Methods for Ambient Air Quality Monitoring, China Environmental Science Press, Beijing, 2006.

1063 Tang, W. Y., Zhao, C. S., Geng, F. H., Peng, L., Zhou, G. Q., Gao, W., Xu, J. M., and Tie, X. X.: Study of ozone  
1064 "weekend effect" in Shanghai, Sci China Ser D, 51, 1354-1360, 10.1007/s11430-008-0088-2, 2008.

1065 Tang, G., Li, X., Wang, Y., Xin, J., and Ren, X.: Surface ozone trend details and interpretations in Beijing,  
1066 2001-2006, Atmos Chem Phys, 9, 8813-8823, 2009.

1067 Tu, J., Xia, Z. G., Wang, H. S., and Li, W. Q.: Temporal variations in surface ozone and its precursors and  
1068 meteorological effects at an urban site in China, *Atmos Res*, 85, 310-337, 10.1016/j.atmosres.2007.02.003,  
1069 2007.

1070 [Wang, H. L., Chen, C. H., Wang, Q., Huang, C., Su, L. Y., Huang, H. Y., Lou, S. R., Zhou, M., Li, L., and Qiao, L.  
1071 P.: Chemical loss of volatile organic compounds and its impact on the source analysis through a two-year  
1072 continuous measurement, \*Atmospheric Environment\*, 80, 488-498, 2013.](#)

1073 Wang, H. X., Zhou, L. J., and Tang, X. Y.: Ozone concentrations in rural regions of the Yangtze Delta in China, *J*  
1074 *Atmos Chem*, 54, 255-265, 10.1007/s10874-006-9024-z, 2006a.

1075 Wang, T. J., Lam, K. S., Xie, M., Wang, X. M., Carmichael, G., and Li, Y. S.: Integrated studies of a  
1076 photochemical smog episode in Hong Kong and regional transport in the Pearl River Delta of China, *Tellus B*,  
1077 58, 31-40, 10.1111/j.1600-0889.2005.00172.x, 2006b.

1078 Wang, X. M., Lin, W. S., Yang, L. M., Deng, R. R., and Lin, H.: A numerical study of influences of urban  
1079 land-use change on ozone distribution over the Pearl River Delta region, China, *Tellus B*, 59, 633-641,  
1080 10.1111/j.1600-0889.2007.00271.x, 2007.

1081 Wang, T., Wei, X. L., Ding, A. J., Poon, C. N., Lam, K. S., Li, Y. S., Chan, L. Y., and Anson, M.: Increasing  
1082 surface ozone concentrations in the background atmosphere of Southern China, 1994-2007, *Atmos Chem Phys*,  
1083 9, 6217-6227, 2009a.

1084 Wang, X. M., Chen, F., Wu, Z. Y., Zhang, M. G., Tewari, M., Guenther, A., and Wiedinmyer, C.: Impacts of  
1085 Weather Conditions Modified by Urban Expansion on Surface Ozone: Comparison between the Pearl River  
1086 Delta and Yangtze River Delta Regions, *Adv Atmos Sci*, 26, 962-972, 10.1007/s00376-009-8001-2, 2009b.

1087 Wei, X. L., Liu, Q., Lam, K. S., and Wang, T. J.: Impact of precursor levels and global warming on peak ozone  
1088 concentration in the Pearl River Delta Region of China, *Adv Atmos Sci*, 29, 635-645,  
1089 10.1007/s00376-011-1167-4, 2012.

1090 Xie, M., Wang, T.J., Jiang, F., Yang, X.Q.: Modeling of natural NOx and VOC emissions and their effects on  
1091 tropospheric photochemistry in China, *Environ. Sci. China*, 28, 31-40, 2007.

1092 Xie, M., Li, S., Jiang, F., and Wang, T. J.: Methane emissions from terrestrial plants over China and their effects  
1093 on methane concentrations in lower troposphere, *Chinese Sci Bull*, 54, 304-310, 10.1007/s11434-008-0402-6,  
1094 2009.

1095 Xie, M., Zhu, K. G., Wang, T. J., Yang, H. M., Zhuang, B. L., Li, S., Li, M. G., Zhu, X. S., and Ouyang, Y.:  
1096 Application of photochemical indicators to evaluate ozone nonlinear chemistry and pollution control  
1097 countermeasure in China, *Atmos Environ*, 99, 466-473, 10.1016/j.atmosenv.2014.10.013, 2014.

1098 Xie, M., Liao, J., Wang, T., Zhu, K., Zhuang, B., Han, Y., Li, M., Li, S.: Modeling of the anthropogenic heat flux  
1099 and its effect on regional meteorology and air quality over the Yangtze River Delta region, China, *Atmos.*  
1100 *Chem. Phys.*, 16, 6071-6089, 10.5194/acp-16-6071-2016, 2016a.

1101 Xie, M., Zhu, K. G., Wang, T. J., Chen, P. L., Han, Y., Li, S., Zhuang, B. L., and Shu, L.: Temporal  
1102 characterization and regional contribution to O<sub>3</sub> and NO<sub>x</sub> at an urban and a suburban site in Nanjing, China,  
1103 *Sci Total Environ*, 551, 533-545, 10.1016/j.scitotenv.2016.02.047, 2016b.

1104 Yarwood, G., Rao, S., Yocke, M., Whitten G.: Updates to the Carbon Bond chemical mechanism: CB05., Final  
1105 Report to the U.S. EPA, RT-0400675, 2005.

1106 Zhang, X. Y., Zhang, P., Zhang, Y., Li, X. J., and Qiu, H.: The trend, seasonal cycle, and sources of tropospheric  
1107 NO<sub>2</sub> over China during 1997-2006 based on satellite measurement, *Sci China Ser D*, 50, 1877-1884,  
1108 10.1007/s11430-007-0141-6, 2007.

1109 Zhang, Y. H., Su, H., Zhong, L. J., Cheng, Y. F., Zeng, L. M., Wang, X. S., Xiang, Y. R., Wang, J. L., Gao, D. F.,  
1110 Shao, M., Fan, S. J., and Liu, S. C.: Regional ozone pollution and observation-based approach for analyzing

1111 ozone-precursor relationship during the PRIDE-PRD2004 campaign, *Atmos Environ*, 42, 6203-6218,  
1112 10.1016/j.atmosenv.2008.05.002, 2008.  
1113 Zhang, Q., Streets, D. G., Carmichael, G. R., He, K. B., Huo, H., Kannari, A., Klimont, Z., Park, I. S., Reddy, S.,  
1114 Fu, J. S., Chen, D., Duan, L., Lei, Y., Wang, L. T., and Yao, Z. L.: Asian emissions in 2006 for the NASA  
1115 INTEX-B mission, *Atmos Chem Phys*, 9, 5131-5153, 2009.  
1116 Zhu, B., Kang, H. Q., Zhu, T., Su, J. F., Hou, X. W., and Gao, J. H.: Impact of Shanghai urban land surface forcing  
1117 on downstream city ozone chemistry, *J Geophys Res-Atmos*, 120, 4340-4351, 10.1002/2014JD022859, 2015.  
1118