

1 Answer to Reviewer #2

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3 1. We corrected "vegetation period" to "growing season" throughout the text.

4

5 2. We agree that better drainage is one of the key factors for establishing of the woody  
6 vegetation. We can illustrate this with shrub colonization along the winter trails near  
7 Nefteugansk (61°55'18" N, 72°36'11" E) (Landsat image), see the attached file. This site is found  
8 in the middle taiga zone so that the soil temperature does not limit the shrub growth. At the  
9 northern treeline however the cause and result puzzle is more complex. Rooting conditions, soil  
10 moisture and water availability, soil thermal conductivity, mineralization and surface heat  
11 balance all found to be significant factors. We rewrite the corresponding paragraph on the page  
12 11 to include three new studies into the discussion, namely, Lloyd et al. (2003), Frost et al.  
13 (2013), and very recent study by Juscak et al. (2016). The conclusion of this discussion supports  
14 our points drawn from area-averaged satellite data analysis. It states that the availability of the  
15 disturbed microsites could be the limiting factor for the woody vegetation proliferation.  
16 Urbanization and subsequent mechanical disturbances and fires create large number of such  
17 disturbances, particularly in 5 km buffer zone, which explains shrubification of this zone.  
18 However, in order to explain higher NDVImax, presence and growth of tall shrubs and trees  
19 are not enough as the field studies suggested. Warmer soils and generally higher air temperature  
20 is needed as it is follows from the rest of the observed literature.

21 3. The manuscript to some degree provides the discussion on the relative roles of climate change  
22 and new infrastructure development (see Introduction, pages 2-3; and Discussion, page 10). In  
23 particular, we use the results of Macias-Fauria et al. (2012) and Bhat et al. (2013). However,  
24 we acknowledge that at present we cannot quantify the relative role of climate versus the  
25 urbanization impact. Therefore, we set to run another study with the aim to attribute the  
26 greening/browning to dominant species in the area.

27

28 **Trends in normalized difference vegetation index (NDVI)**  
29 **associated with urban development in northern West Siberia**

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## 11 **Abstract**

12 Exploration and exploitation of oil and gas reserves of northern West Siberia has promoted  
13 rapid industrialization and urban development in the region. This development leaves  
14 significant footprints on the sensitive northern environment, which is already stressed by the  
15 global warming. This study reports the region-wide changes in the vegetation cover as well as  
16 the corresponding changes in and around 28 selected urbanized areas. The study utilizes the  
17 normalized difference vegetation index (NDVI) from high-resolution (250 m) MODIS data  
18 acquired for summer months (June through August) during 15 years (2000–2014). The results  
19 reveal the increase of NDVI (or “greening”) over the northern (tundra and tundra-forest) part  
20 of the region. Simultaneously, the southern, forested part shows the widespread decrease of  
21 NDVI (or “browning”). These region-wide patterns are, however, highly fragmented. The  
22 statistically significant NDVI trends occupy only a small fraction of the region. Urbanization  
23 destroys the vegetation cover within the developed areas and at about 5–10 km distance around  
24 them. The studied urbanized areas have the NDVI values by 15% to 45% lower than the  
25 corresponding areas at 20–40 km distance. The largest NDVI reduction is typical for the newly  
26 developed areas, whereas the older areas show recovery of the vegetation cover. The study  
27 reveals a robust indication of the accelerated greening near the older urban areas. Many Siberian  
28 cities become greener even against the wider browning trends at their background. Literature  
29 discussion suggests that the observed urban greening could be associated not only with special

1 tending of the within-city green areas but also with the urban heat islands and succession of  
2 more productive shrub and tree species growing on warmer sandy soils.

3

## 4 **1 Introduction**

5 Significant shifts in the vegetation land cover and biological productivity manifest rapid climate  
6 change in the northern high latitudes (Hinzman et al., 2005; Groisman and Gutman, 2013). As  
7 in other polar regions, biomes of northern West Siberia (Fig. 1), hereafter referred to as NWS,  
8 respond on global warming with complex patterns of multidirectional changes (Bunn and  
9 Goetz, 2006; Lloyd and Bunn, 2007; McDonald et al., 2008; Walker et al., 2009; Bhatt et al.,  
10 2013). Although the patterns are geographically fragmented (Elsakov and Teljatnikov, 2013;  
11 Barichivich et al., 2014; Guay et al., 2014), the northern biomes (tundra and forest-tundra)  
12 demonstrate widespread increase of the vegetation productivity (“greening”). Its attribution to  
13 tall shrub and graminoids in tundra ecosystems (Frost and Epstein, 2014) and enhanced tree  
14 growth in forest-tundra ecosystems (Urban et al., 2014) suggested that we observe transitions  
15 to alternative, potentially more productive ecosystems rather than anomalously enhanced  
16 biological productivity in response to global warming (Kumpula et al., 2012; Macias-Fauria et  
17 al., 2012). These environmental shifts have been unfolding for at least the last three decades  
18 (Keeling et al., 1996; Myneni et al., 1997) albeit at the slower pace since 2003 (Bhatt et al.,  
19 2013).

20 Simultaneously, the southern biomes (northern and middle taiga forest to the south of  
21 about 65°N) demonstrate decrease of productivity (“browning”). This browning has been  
22 robustly associated with decreasing mass of the green leaves, which has been linked to Siberian  
23 forest acclimation without the transition to an alternative ecosystem (Lapenis et al., 2005; Lloyd  
24 and Bunn 2007). Although the grown trees are stressed by increasing temperatures, other types  
25 of the vegetation cover (grass- and shrub-lands, mires) and disturbed patches with young tree  
26 seedlings exhibit greening within the same bioclimatic zone.

27 The persistent greening trends of the disturbed vegetation cover is in focus of this study.  
28 Over the last 30–50 years, exploration of oil and gas reserves promoted rapid industrialization  
29 and urban development of NWS. It makes this part of the circumpolar region particularly  
30 interesting for studies of the emerging alternative ecosystems. The development has left  
31 significant footprints on the vegetation cover. These footprints were found not only at the  
32 development sites but also across large distances (Walker et al., 2009; Kumpula et al., 2012).

1 More specifically, Kumpula et al. (2011) reported extensive transformation from shrub- to more  
2 productive (greener) grass- and sedge-dominated tundra in NWS that reclaim artificial terrain  
3 disturbances at Bovanenkovskiy gas field, Yamal peninsula (70°N). Similar environmental  
4 shifts towards more productive plant successions on disturbed patches were found in studies of  
5 post-mined sand and sandy loam quarries around Noyabrsk at 63°N (Koronatova & Milyaeva,  
6 2011). This transformation across open woodlands and forest-tundra was found to be more  
7 ambiguous. There has been found 26% (from 2316 to 1715 g m<sup>-2</sup>) decrease of the total  
8 aboveground biomass due to enhanced bog formation on the background of widespread  
9 greening of better drained sandy patches (Moskalenko, 2013). Detailed study by Sorokina  
10 (2013) revealed the increased soil temperature by 1K to 4K on the disturbed patches within  
11 forest, mire and tundra ecosystems. The reclaiming vegetation cover is characterized by  
12 increasing share of shrubs (from 18.5% to 27.8% in the forest ecosystems) and grasses (from  
13 less than 0.1% to 8.8%) in its above-ground biomass.

14 Although disturbed patches occupy a relatively small fraction in NWS, their  
15 contribution to the observed vegetation productivity changes can be disproportionately  
16 significant. For example, for the Yamal peninsula, Elsakov and Teljatnikov (2013) concluded  
17 that its widely referred greening (Macias-Fauria et al., 2012) could be traced to statistically  
18 significant changes over just 4.8% of the total area. Moreover, this greening was linked to shrub  
19 growth on the patches with disturbed permafrost or within the patches of already establish  
20 shrubs (Walker et al., 2009). Barichivich et al. (2014) noted that browning in southern forest  
21 biomes could be also partially associated with disturbances.

22 The reviewed literature broadly agrees that the observed transformation of vegetation  
23 cover represents an alternative ecosystem state (Kumpula et al., 2012; Macias-Fauria et al.,  
24 2012) and a new geocological regime (Raynolds et al., 2014) acclimated to higher  
25 temperatures that are likely to persist. This point of view received strong support from the plot-  
26 scale International Arctic Tundra Experiment, which simulated warmer climate conditions in  
27 high latitudes (Walker et al., 2006; Elmendorf et al., 2012). The question remains whether the  
28 plot-scale experiments and field studies at a few carefully selected locations, e.g., Vaskiny  
29 Dachi in the central Yamal peninsula (Leibman et al., 2015), could be extrapolated for  
30 assessment of the larger scale transition to alternative ecosystems. To approach this question,  
31 we use an opportunity provided by the above-mentioned intensive urbanization in NWS. The  
32 urbanization created large number of relatively large-scale (tens of km across) artificial

1 disturbances across all major regional biomes. This study looks at systematic differences in the  
2 apparent vegetation productivity trends between the undisturbed background and the areas  
3 disturbed by urbanization and industrial development.

4 As in several previously published studies, this study utilizes a Normalized Difference  
5 Vegetation Index (NDVI) derived from satellite data products from the Moderate Resolution  
6 Imaging Spectroradiometer (MODIS) onboard of the Earth Observing System-Terra platform  
7 satellite. Studies by Frey and Smith (2007), Epstein et al. (2012), Macias-Fauria et al. (2012)  
8 and others demonstrated that MODIS NDVI could be successfully used in vegetation  
9 productivity assessment for Siberian biomes. This study details the previously reported NDVI  
10 maps of the region and expands them over the last 15 years (2000–2014) period. The vegetation  
11 productivity of emerging alternative ecosystems are calculated at and around of 28 urbanized  
12 areas, hereafter referred to as cities. These 28 cities are located in four northern biomes ranging  
13 from tundra (Bovanenkovo) to middle taiga forest (Uray).

14 This paper has the following structure: Data and methods of the study are described in  
15 Section 2. Section 3 presents the analysis and results for both the background changes in NDVI  
16 and NDVI in and around the selected cities. Section 4 discusses the results in the context on  
17 previous research and emphasize the role of the cities. Section 5 outlines conclusions.

## 18 **2 Data and methods**

19 This study is based on analysis of the maximum summer NDVI, which we denote as  $NDVImax$ .  
20  $NDVImax$  was obtained from MODIS NDVI 16-day composites with the 250-m spatial  
21 resolution (MOD13Q1). The data were downloaded from the NASA’s Earth Observing System  
22 Data and Information System (EOSDIS) for 15 summers (June–July–August, JJA) covering the  
23 period 2000–2014.

### 24 **2.1 The maximum Normalized Difference Vegetation Index, $NDVImax$**

25 NDVI is defined as a normalized ratio of reflectance factors in the near infrared (NIR) and red  
26 spectral radiation bands

$$27 \quad NDVI = \frac{\rho_{NIR} - \rho_{RED}}{\rho_{NIR} + \rho_{RED}} \quad (1)$$

28 where  $\rho_{NIR}$  and  $\rho_{RED}$  are the surface bidirectional reflectance factors for their respective  
29 MODIS bands on the TERRA platform satellite. NDVI exploits the contrast between the red  
30 and NIR reflectance of vegetation, as chlorophyll is a strong absorber of the red light, while the

1 internal structure of leaves reflects highly in the NIR. The greater the difference between the  
2 reflectance in the red and NIR portions of the spectrum, the more chlorophyll is found in  
3 vegetation canopy. Vegetation generally yields positive NDVI values, which approach +1 with  
4 increasing plant chlorophyll content or green aboveground biomass. NDVI with the values  
5 below 0.2 generally corresponds to non-vegetated surfaces, whereas green vegetation canopies  
6 have NDVI greater than 0.3.

7 NDVI in remote sensing studies is a popular proxy for gross photosynthesis, and  
8 therefore, for vegetation productivity (Goetz et al., 2005). A strength of NDVI is its  
9 normalization, which makes it relatively insensitive to radiometric attenuation (e.g., by cloud  
10 shadows) present in multiple bands. The main weakness of NDVI is its inherent nonlinearity  
11 that leads to asymptotic saturation of NDVI over higher biomass conditions. This saturation,  
12 also known as the NDVI degradation, is particularly strong in the areas with higher canopy  
13 background brightness corresponding to the most productive biomes. By contrast, the biomass  
14 accumulation in the less productive biomes could be approximated with a linear regression  
15 model. NDVI typically does not degrade in tundra (Raynolds et al., 2012), but the degradation  
16 should be empirically taken into account in NDVI interpretations for more southern biomes  
17 (D'Arrigo et al., 2000; Zhang et al., 2004).

18 A remarkably strong correlation ( $R^2 = 0.94$ ,  $p < 0.001$ ) was found between total above-  
19 ground phytomass sampled at the peak of summer and the maximum annual NDVI (*NDVImax*)  
20 in studies of the North America and Eurasia transects (Raynolds et al., 2012). This strong  
21 relationship encouraged us to use *NDVImax* in this study. *NDVImax* is a more conservative  
22 characteristic of the vegetation cover that is linked to the total biomass at the late phenological  
23 phases. Moreover, *NDVImax* eliminates seasonal variations and the relative shifts between  
24 phenological phases in different climatic zones. Hence, *NDVImax* is particularly convenient for  
25 and frequently used in the environmental studies dealing with long-term and large-scale  
26 changes, including effects of the climate change.

27 *NDVImax* was obtained from MOD13Q1 data product. This product is distributed in  
28 adjacent non-overlapping tiles with the side of approximately 10 degrees (at the equator) and  
29 the Sinusoidal (SIN) tile grid projection (Solano et. al, 2010). Five tiles to cover the entire area  
30 of interest (total 20 tiles per each summer) were downloaded and imported into the ArcGIS  
31 geographic information system. Images were combined and re-projected from the original SIN  
32 to the Universal Transverse Mercator projection (UTM Zone 42N, WGS84 ellipsoid). The data

1 were quality-filtered by the MODIS reliability data provided together with the MOD13Q1  
 2 product. Only data of the highest quality, which excluded snow/ice- and cloud covered pixels,  
 3 were retained. The  $NDVI > 0.3$  criterion was used here to exclude water, bare soil and other  
 4 non-vegetated pixels from the analysis. The data gaps in the raster mosaic pixels were then  
 5 filled with information using the nearest neighbor statistical interpolation from the surrounding  
 6 pixels with data. Finally,  $NDVImax$  maps for each summer were obtained through identification  
 7 of the maximum NDVI value from each 16-day composite for each pixel. The analysis operates  
 8 with these  $NDVImax$  maps of the 250-m resolution covering the 15-year period 2000–2014.

## 9 **2.2 $NDVImax$ in and around selected cities**

10 The NWS territory provides a unique opportunity for statistical study, which compares the  
 11 effect of recent urbanization along with the effect on the climate change on the vegetation land  
 12 cover in and around cities across several northern biomes. Twenty-eight (28) cities, some of  
 13 them with more than 100,000 inhabitants (see Table 1), were selected.  $NDVImax$  was studied  
 14 within 40 km buffer zones around each city. Each buffer zone was broken into 8 rings of 5 km  
 15 width centered at the city core zone. This approach is similar to one used by Zhang et al. (2004).

16 Let us introduce  $\langle NDVImax(t, i, n) \rangle$  as the  $NDVImax$  value in the year  $t = 1 \dots 15$  for  
 17 the ring number  $i = 0 \dots 8$  (the ring  $i = 0$  corresponds to the city core and  $i = 8$  – to the most  
 18 distant background ring) in the city  $n = 1 \dots 28$  that is averaged over all pixels within the ring.  
 19 This approach does not differentiate between the vegetation productivity changes and the  
 20 changes of the vegetation biological composition (succession). In this sense, it will be  
 21 insensitive to the  $NDVImax$  trends due to urban expansion, shifting vegetation disturbances and  
 22 re-vegetation. As the  $\langle NDVImax(t, i, n) \rangle$  values strongly vary from city to city and across  
 23 different biomes, it is convenient to introduce a relative footprint of a city as

$$24 \quad F(t, i, n) = \frac{\langle NDVImax(t, i, n) \rangle}{\langle NDVImax(t, n) \rangle} \quad (2)$$

25 and the relative linear trends,  $R(i, n)$ , which are computed by the least-squares fit of  $F(t, i, n)$   
 26 to the first-order polynomial for the ring  $i$  in the city  $n$ ,  $\overline{\langle NDVImax(t, n) \rangle}$  is time-averaged  
 27  $\langle NDVImax(t, i, n) \rangle$ . Because the study operates with rather short 15-year time series, the years  
 28 with the maximum and the minimum  $NDVImax$  can strongly impact the trends. These two years  
 29 (2002 and 2014 for the majority of cities) can be considered as outliers as their  $NDVImax$  were  
 30 beyond three standard deviations of the respective time series. We demonstrate this for four

1 cities (Bovanenkovsky, Nadym, Noyabrsk and Surgut) in Fig. 3. Therefore, the years with  
2 minimum and maximum  $NDVImax$  were excluded from the trend fitting. We will also consider  
3 the differences in the relative footprints,  $\Delta F(i, j, n) = \overline{F(t, i, n)} - \overline{F(t, j, n)}$ , and relative  
4 differences in the trends,  $\Delta R(i, j) = R(i, n) - R(j, n)$ , between the rings  $i$  and  $j$  for the city  $n$ .  
5 The urbanization footprint could be characterized through the divergent trends:  $\Delta R(0,8)$  –  
6 between the city core and the corresponding natural land cover;  $\Delta R(0,5)$  – between the core  
7 and the first ring around it; and  $\Delta R(5,8)$  – between supposedly the most affected 5-km ring and  
8 the background 40-km ring.

### 9 **3 Results**

#### 10 **3.1 Regional $NDVImax$ patterns and trends**

11 Our study expands, updates and details the results of the  $NDVImax$  analysis for 1981–1999 by  
12 Zhou et al. (2001), 1982–2003 by Bunn and Goetz (2006), 1982–2008 by Beck and Goetz  
13 (2011), 1982–2011 by Barichivich et al. (2014) and 2000–2009 by Elsakov and Teljatnikov  
14 (2013). Two novel aspects should be mentioned in this context: (1) Whereas the previous  
15 studies analyzed coarse-resolution data, which is likely to exaggerate the extent and magnitude  
16 of the  $NDVImax$  trends (Zhao et al., 2009; Elsakov and Teljatnikov, 2013), we use the fine-  
17 resolution (250 m) data; and (2) Fine-resolution data give an opportunity to trace the changes  
18 to specific biomes within the same bioclimatic zone and to reveal effects of urban disturbances.

19 The updated mean  $NDVImax$  and  $NDVImax$  trend maps (2000–2014) are shown in Fig.  
20 1a and 1b, respectively.  $NDVImax$  in NWS generally decreases from the southwest to the  
21 northeast of the territory. The largest  $NDVImax$  values (the most productive vegetation) are  
22 found along the Ob river and between the Ob river and Ural mountains, whereas the central,  
23 swamped part of NWS has much lower  $NDVImax$ . We observe that  $NDVImax$  is significantly  
24 higher on river terraces with better-drained, sandy soils, which are warmer in summertime and  
25 have deeper seasonal active layer. The new maps confirm continuing widespread greening in  
26 tundra and forest-tundra biomes. However, this greening is highly fragmented and to the large  
27 degree could be associated with sandy soils as well as with smaller greening patches associated  
28 with permafrost destruction, landslides, thermokarst and other local disturbances. Fig. 1b shows  
29 the statistically significant trends. It clearly demonstrates that the previously reported  
30 widespread greening trends are statistically insignificant. The most significant areas of greening  
31 are found in Taz and southern Gydan peninsulas. This finding is in good agreement with



1 previously reported plot scale studies using LANDSAT images. Table 2 and Fig. 2 aggregate  
2 the greening and browning trends in the NWS biomes. They show that the forest biomes exhibit  
3 more widespread and larger magnitude browning. The maximum area fraction of browning  
4 (21.3%) was found in middle taiga and the minimum area fraction (8.9%) – in tundra biomes.  
5 Contrary, the area fraction of greening is the largest (81.7%) in forest-tundra and the smallest  
6 (35.5%) in middle taiga biomes.

7         Comparison with Lloyd and Bunn (2007), Bhatt et al. (2013) and Elsakov and  
8 Teljatnikov (2013) studies further reveals that the area of more productive vegetation cover  
9 continue to grow as well as the production in shrub- and graminoid-dominated ecosystems.  
10 Despite colder recent winters (Cohen et al., 2013) and somewhat damped summer warming  
11 (Tang and Leng, 2012), the greening now dominates the changes in all four biomes. We observe  
12 the strongest greening near the southern tundra boundary. This pattern is consistent with  
13 previously noted shrubification and treeline advance in this area (Devi et al., 2008; MacDonald  
14 et al., 2008).

### 15 **3.2 *NDVImax* patterns and trends in and around 28 cities**

16 Disturbances of the vegetation cover around 28 cities in NWS considerably modify the  
17 observed complex pattern of the background *NDVImax* trends. In this analysis, we distinguish  
18 the city core,  $i = 0$ , with strongly disturbed vegetation cover and therefore low *NDVImax* and  
19 the rings  $i = 1 \dots 8$  where the area of disturbances progressively decrease with the distance from  
20 the city core. Fig. 3 shows the analysis of *NDVImax* changes for four typical cities located in  
21 four different biomes. The observed convergence of the statistical properties of *NDVImax* in  
22 the rings  $i = 4 \dots 7$  to those in the ring 8 supports the intuitive assumption that the area fraction  
23 of urban disturbances is gradually reduced with the distance from the city. Indeed, the  
24 correlation coefficients between time series of  $F(t, 0, n)$  and  $F(t, i, n), i = 1 \dots 8$ , are  
25 decreasing  $i$  (see the upper panels in Fig. 3).

26         Now, we consider differences between the background *NDVImax* trends and the trends  
27 over the disturbed vegetation cover in and around the cities. Fig. 4a and 4c show that all cities  
28 have strongly reduced *NDVImax* values. Some cities have *NDVImax* reduced by more than 30%  
29 as compared to the background. As it has been already suggested by Fig. 3 (for Nadym and,  
30 particularly, for Noyabrsk), the closest 5-km ring ( $i = 1$ ) often exhibits higher *NDVImax* than  
31 the more distant background. This unusual feature could be traced down to the preferential

1 location of the cities on generally greener river terraces. So that the greener patches contribute  
2 more heavily to the mean *NDVImax* of the inner rings.

3         Apart from lower *NDVImax* in the city cores, Fig. 4d reveals two opposite dependences  
4 between the vegetation disturbance and the vegetation productivity. The cities with larger  
5 relative *NDVImax* reduction (the large negative  $\Delta F$ ) demonstrate the accelerated *NDVImax*  
6 recovery (the large positive  $\Delta R$ ). This rapid recovery at the initial stages of the vegetation  
7 succession has been repeatedly noted in several plot-scale studies (Sorokina, 2003; Archegova,  
8 2007; Koronatova and Milyaeva, 2011) and high-resolution satellite image analyses (Kornienko  
9 and Yakubson, 2011; Kumpula 2011). Unfortunately, majority of these studies have been  
10 published in Russian only. Intercomparison with Fig. 4c, where this dependence is less  
11 pronounced, suggests that the vegetation cover in cities is more resistant to the stress in the  
12 extreme warm and cold years. Moreover, the protecting effect of the city is stronger for the  
13 northern cities responding with strong greening due to extremely warm summer temperatures.

14         It is interesting to observe that the effect of cities does not exhibit clear regular  
15 dependence on the city population or specific location. The only visible effect in Fig. 4a could  
16 be attributed to the age of the city – the younger northern cities continue to destroy vegetation  
17 cover as they expand, whereas the vegetation cover in the established southern cities recovers.  
18 Fig 4b shows that the disturbances have the largest positive effect on both the mean *NDVImax*  
19 and its trends in the forest-tundra biome. Here, 6 out of 7 cities induce the accelerated greening  
20 in the 5-km ring.

21         The biome-averaged impact of the urban disturbances on *NDVImax* is given in Fig. 5.  
22 The analysis show that the city footprint is visible at large distances. Even at the 15 km distance  
23 from the city core, *NDVImax* is systematically higher in the northern biomes and lower in the  
24 middle taiga biome. Moreover, *NDVImax* is also systematically the highest in the closest 5-km  
25 ring where the area of disturbances is the largest. Fig. 6 illustrates the vegetation changes  
26 leading to higher biological productivity of the disturbed land patches in the northern biomes.  
27 The newly established tree seedlings and shrubs, which are more productive relative to the  
28 background in the northern biomes, are less productive relative the background of the mature  
29 trees in the middle taiga biome.

## 1 4 Discussion

2 The results of the presented high-resolution *NDVImax* analysis in 2000–2014 show continuing  
3 changes in the vegetation land cover and vegetation productivity in NWS. In general, the  
4 northern biomes (tundra and forest-tundra north of about 65°N) and to some degree open  
5 swamped areas within the forest biomes demonstrate persistent greening, whereas the forested  
6 areas showed similarly persistent browning. More detailed analysis has however revealed that  
7 the statistically significant changes are highly fragmented. Such changes occupy only a minor  
8 fraction of NWS and often could be collocated with natural or anthropogenic disturbances of  
9 the vegetation cover. This observation lead Macias-Fauria et al. (2012) and Kumpula et al.  
10 (2012) to hypothesize that those disturbances help to establish alternative, more productive  
11 ecosystems of graminoids, shrubs and tree seedlings. In this discussion, we consider the major  
12 climatological and physical factors that may support the alternative vegetation cover.

13 The cold continental climates of the NWS (the Köppen–Geiger climate types Dfc and  
14 ET) with relatively high amount of annual precipitation determine a short **vegetation**  
15 **periodgrowing season** and strong dependence of the maximum vegetation productivity on  
16 summer temperatures (Barichivich et al., 2014). This dependence is not of universal character  
17 across biomes. The higher spring–summer temperatures (Bhatt et al., 2013; Ippolitov et al.,  
18 2014) and larger amount of accumulated snow (Bulygina et al., 2014) favor the tall deciduous  
19 shrub (*Betula nana* and different *Salix species*) in tundra and forest-tundra (e.g Sturm et al.,  
20 2001; Elmendorf et al., 2012). Similarly, grasses (graminoids) respond on higher spring  
21 temperatures with higher biomass production. It should be emphasized however that the  
22 associations between the *NDVImax* trends and productivity of the specific ecosystems within  
23 the arctic biomes remain rather loose (Frey and Smith, 2007; Kornienko and Yakubson, 2011).  
24 Moreover, even synthetic *NDVImax* time series (Guay, et al., 2014) are too short and too  
25 variable for statistically robust conclusions over the major part of the area.

26 The response on the observed climate change is different in the forest biomes.  
27 Summertime surface air temperatures increased only weakly in central and southern NWS  
28 (Ippolitov et al., 2014) likely being damped by increasing cloud cover (Tang and Leng, 2012;  
29 Esau et al., 2012). The temperature trends in the winter months were negative (Cohen et al.,  
30 2013; Outten and Esau, 2011; Outten et al., 2013). Thus, there were no consistent warming  
31 trends over this territory but rather a few summer heat waves (e.g. in 2002, 2007, 2012) with  
32 certain impact on the forest productivity.

1 As the reviewed literature disclose, the direct impact of the climate factors on the  
2 vegetation cannot be established unambiguously. Even conclusions derived from the analysis  
3 of common data sets differ on the relative role of the climate change and the natural decadal  
4 variability. For example, Melnikov et al. (2004) concluded that the climatic trends are only  
5 weakly discernable in the active soil layer data, whereas Streletsy et al. (2012) found a  
6 significant increase (up to 0.3 m) of the active layer thickness. Moreover, simulations of the  
7 forest biomass productivity with a stand-alone dynamic vegetation model by Schumann and  
8 Shugart (2009) produced no significant productivity change at Siberian forest sites for the  
9 warming below 2K. At the same time, the observed biome-wide dichotomy of the grass-covered  
10 area greening versus the tree-covered area browning was reproduced in the anthropogenic  
11 global warming experiments with the Community Atmospheric Model version 3 with the Lund-  
12 Potsdam-Jena dynamic vegetation model (Jeong et al., 2012).

13 A more coherent picture has been suggested by studies of vegetation reclaiming the  
14 natural or anthropogenic disturbances (Kornienko and Yakubson, 2011; Kumpula, 2011). Lloyd  
15 et al (2003) concluded that the establishment and seasonal growth of woody species at the  
16 Alaska treeline was likely limited by the availability of disturbed well-drained microsites. They  
17 suggested that the vegetation response to regional climatic changes is presently controlled by  
18 the availability of such microsites. This study supports the idea of the changes limited to the  
19 disturbed microsites availability. We found~~The identified~~, very fragmented pattern of changes  
20 in the NWS. and~~The~~ larger changes near the cities are consistent with~~suggest~~ the large impact  
21 of re-populated microsites with stronger woody vegetation on~~of soil and vegetation cover~~  
22 disturbances on~~of~~ the *NDVImax* trends. In this sense, human activity and urbanization create  
23 variety of the disturbed microsites needed for establishing of more productive vegetation.

24 The soil disturbances significantly modify both the surface heat balance and the  
25 drainage conditions in the area. Pavlov and Moskalenko (2002) and more recently Yu et al.  
26 (2015) showed that the disturbed soils in tundra and northern taiga (near Nadym) are warmer  
27 and accumulated more heat (thaw deeper) during the summer seasons. Establishing of woody  
28 species may revert this process. Thus, although -Lloyd et al. (2003) found that woody species  
29 are more readily establishing on disturbed microsites with better drainage and drier warmer  
30 soils, Juscak et al. (2016) showed poor soil thermal conductivity below grown dwarf shrub  
31 microsites, which results in thinner active soil layer and worse rooting conditions. Frost et al.  
32 (2013) studies in our NWS area strongly support the envisaged transformation of the soil

1 conditions. They reported higher soil temperatures during the colonization phase and lower  
2 temperatures during the mature phase of *Alnus* growth. Fig. 6 and 7 illustrate the shrub  
3 establishment on disturbed microsites, ~~and that the disturbed soils are better drained, so that the~~  
4 ~~latent heat flux is reduced. More general discussion however should account for other physical~~  
5 ~~and biological processes in the active soil layers. Better drainage gives better rooting conditions~~  
6 ~~for forbs and trees as well as increases organic mass loss in soils, litter decomposition and~~  
7 ~~decreases moss productivity (Hicks Pries et al., 2013). Yakubson et al. (2012) reported high~~  
8 ~~correlation between the enhanced *NDVImax* and dryer soils at the disturbed areas around~~  
9 ~~Bovanenkovo. Similar dependences were found over Taz peninsula (Kornienko and Yakubson,~~  
10 ~~2011).~~

11 Mosaic of vegetation changes on “microsites” is readily observable in field studies.  
12 More general discussion with relevance to the observed changes in the biomass production,  
13 however, should account for other physical and biological processes in the active soil layers.  
14 Better drainage gives better rooting conditions for forbs and trees as well as increases organic  
15 mass loss in soils, litter decomposition and decreases moss productivity (Hicks Pries et al.,  
16 2013). Yakubson et al. (2012) reported high correlation between the enhanced *NDVImax* and  
17 dryer soils at the disturbed areas around Bovanenkovo. Similar dependences were found over  
18 Taz peninsula (Kornienko and Yakubson, 2011).

19 ~~Moreover~~, Brunzel et al. (2011) showed that the boundary layer dynamics with enhanced micro-  
20 circulations over the microsites might additionally increase the surface temperature  
21 heterogeneity, enhancing the sensible heat flux by as much as  $50 \text{ Wm}^{-2}$  for larger disturbances.  
22 Our results for the 5-km ring seem to support this feedback hypothesis.

23 Summarizing this discussion on the relative role of the changes in drainage and the  
24 surface heat balance, we comment that while woody vegetation demonstrates a clear  
25 phenotypical response on better drainage conditions, the area-average biomass productivity has  
26 stronger response on the soil temperature as found by Barichivich et al. (2014).

27 Cities and industrial installations in NWS are frequently built-up on sand beds. It may  
28 partially explain the more positive *NDVImax* trends found in and around the city cores.  
29 Koronatova and Milyaeva (2011) studied plant succession over 1999–2010 in post-mined sandy  
30 and loam quarries around Noyabrsk city. The quarries were colonized by pine seedlings already  
31 during the first 5 years, skipping the grass community stage, and an open pine community with  
32 green moss and lichen species was established by year 20.

1           Finally, it is fruitful to comment on the relative roles of climate changes and the  
2 anthropogenic disturbances introduced through the new infrastructure development. Yu et al.  
3 (2015) compared the vegetation changes around Nadym using high spatial resolution imagery  
4 acquired in 1968 and 2006. This area corresponds to the buffer zones shown in Fig. 3 for this  
5 city. They concluded that about 9% of the area revealed increase in vegetation cover in response  
6 to climate warming while 10.8% of the area had decrease in vegetation cover due to the  
7 infrastructure development and related factors (logging, tracking etc). The direct mechanical  
8 impact on the vegetation cover was very localized (mostly within 100 m from the infrastructural  
9 objects), but indirect biophysical impacts, such as changes in the surface hydrology, heat  
10 balance and ecosystem damages (e.g. fires) were found over significantly larger areas. These  
11 wider indirect impacts of urbanization are visible not only in highly aggregated analysis in the  
12 present study but also in plot-scale experiments. The global warming favors expansion of wood  
13 vegetation over the permafrost. However, the mechanical removal of shrubs in a Siberian tundra  
14 site initiated permafrost thaw converting the plot into waterlogged depression within five years  
15 (Nauta et al., 2014). Thus, there seems to be a concert between the impacts on vegetation cover  
16 and the soil cryosphere induced by global warming and anthropogenic disturbances.

17

## 18 **5 Conclusions**

19 This study presented the *NDVImax* analysis of the MODIS NDVI data product MOD13Q1 with  
20 250 m spatial resolution over 2000–2014. We obtained the maps of *NDVImax* and *NDVImax*  
21 trends for the northern West Siberia region where intensive oil and gas exploration created a  
22 large number of anthropogenic disturbances. These new maps were discussed in the context of  
23 the previously published low- and high-resolution NDVI analysis available for this region. The  
24 obtained *NDVImax* trends are highly fragmented, but confirm the observed dichotomy of  
25 northern greening (the increasing vegetation productivity) versus southern browning (the  
26 decreasing productivity). The statistically significant *NDVImax* trends occupy only a small  
27 fraction of NWS. The most significant trends are found on the territories with sandy soils and  
28 with larger concentration of soil disturbances. Thus, the new map substantially corrects the  
29 previous picture of the vegetation cover changes in NWS and suggests stronger resilience of  
30 undisturbed vegetation cover to the climate change on decadal time scales.

31           It has been proposed that the disturbances might help to establish alternative, more  
32 productive vegetation cover. We used *NDVImax* data in and around 28 regional urbanized and

1 industrial areas to compare the vegetation productivity of the reclaiming and background plant  
2 communities across all four major biomes. We assumed growing concentration of disturbances  
3 towards the city cores. The results indicated that the reclaiming plant communities tend to be  
4 more productive. Only some cities in middle taiga biome, notably the largest cities of Surgut  
5 and Nizneartovsk on the Ob river, exhibit lower productivity in the closest city proximity.

6 As it was expected, we found  $NDVImax$  at the city cores to be 15% to 40% lower than  
7 over the corresponding background. At the same time, many cities has become significantly  
8 greener over the analyzed 15 years. This tendency reflects both the targeted efforts to create  
9 more environmental friendly residential areas and the urban heat island impact on the active  
10 soil layer thickness and drainage.

11

## 12 **Appendix A: Statistical significance of inter-city analysis**

13 The  $NDVImax$  data in and around the cities are very variable. In order to estimate the statistical  
14 significance of the difference of multiple time series, we used a conservative hypothesis testing  
15 with the Student's  $t$ -criterion. In this analysis, two statistical parameters are different at the  
16 confidence level 95% ( $\alpha = 1.96$ , the light gray shading in Figs. 3 and 4) and 99% ( $\alpha = 2.58$ ,  
17 the dark gray shading) assuming the normal distribution of the difference errors when

$$18 \sqrt{\frac{2}{N-1} \max_{m=1\dots 28} \left( \max_{k=0\dots 8} \sigma_n(k) \right)} > \alpha \quad (A1)$$

19 where  $\sigma_n(k)$  is the standard deviation of a parameter in testing for the ring  $k$  and the city  $n$ . We  
20 want to stress that the conservative estimation is related to the differences, which have  
21 significantly larger magnitude than the variations of  $NDVImax$  within each ring. Thus, the  
22 differences appear to be more significant than the trends for each of the rings. There were found  
23 no significant trends at levels higher than 90% for any 5-km ring around any city.

24

25

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5



## 1 **References**

- 2 Archegova, I. B.: Thermal regime of tundra soils under reclamation and restoration of natural  
3 vegetation, *Eurasian Soil Science*, 40, 854–859, 2007.
- 4 d'Arrigo, R. D., Malmstrom, C. M., Jacoby, G. C., Los, S. O., and Bunker, D. E.: Correlation  
5 between maximum latewood density of annual tree rings and NDVI based estimates of forest  
6 productivity, *Int. J. Remote Sens.*, 21, 2329–2336, 2000.
- 7 Barichivich, J., Briffa, K. R., Myneni, R., van der Schrier, G., Dorigo, W., Tucker, C. J.,  
8 Osborn, T. J., and Melvin, T. M.: Temperature and snow-mediated moisture controls of summer  
9 photosynthetic activity in northern terrestrial ecosystems between 1982 and 2011, *Remote*  
10 *Sens.*, 6, 1390–1431, 2014.
- 11 Beck, P., and Goetz, S.: Satellite observations of high northern latitude vegetation productivity  
12 changes between 1982 and 2008: ecological variability and regional differences, *Environ. Res.*  
13 *Lett.*, 6, 045501, 2011.
- 14 Bhatt, U. S., Walker, D. A., Raynolds, M. K., Bieniek, P. A., Epstein, H. E., Comiso, J. C.,  
15 Pinzon, J. E., Tucker, C. J., and Polyakov, I. V.: Recent declines in warming and vegetation  
16 greening trends over pan-Arctic tundra, *Remote Sens.*, 5, 4229–4254, 2013.
- 17 Brunsell, N. A., Mechem, D. B., and Anderson, M. C.: Surface heterogeneity impacts on  
18 boundary layer dynamics via energy balance partitioning, *Atmos. Chem. Phys.*, 11, 3403–3416,  
19 2011.
- 20 Bulygina, O., Groisman, P. Y., Razuvaev, V., and Korshunova, N.: Changes in snow cover  
21 characteristics over Northern Eurasia since 1966, *Environ. Res. Lett.*, 6, 045204, 2011.
- 22 Bunn, A., and Goetz, S.: Trends in Satellite-Observed Circumpolar Photosynthetic Activity  
23 from 1982 to 2003: The influence of seasonality, cover type, and vegetation density, *Earth*  
24 *Interactions*, 10, 1–19, 2006.
- 25 Cohen, J. L., Furtado, J. C., Barlow, M. A., Alexeev, V. A., and Cherry, J. E.: Arctic warming,  
26 increasing snow cover and widespread boreal winter cooling, *Environ. Res. Lett.*, 7, 014007,  
27 2012.
- 28 Devi, N., Hagedorn, F., Moiseev, P., Bugmann, H., Shiyatov, S., Mazepa, V., and Rigling, R.:  
29 Expanding forests and changing growth forms of Siberian larch at the Polar Urals treeline  
30 during the 20th century, *Global Change Biology*, 14, 1581–1591, 2008.

- 1 Elmendorf, S., Henry, G., Hollister, R., et al.: Plot-scale evidence of tundra vegetation change  
2 and links to recent summer warming, *Nature Climate Change*, 2, 453–457, 2012.
- 3 Elsakov, V., and Teljatnikov, M.: Effects of interannual climatic fluctuations of the last decade  
4 on NDVI in north-eastern European Russia and Western Siberia, *Contemporary Problems of*  
5 *the Earth's Remote Sensing*, 10(3), 260–271, 2013.
- 6 Epstein, H. E., Reynolds, M. K., Walker, D. A., Bhatt, U. S., Tucker, C. J., and Pinzon, J. E.:  
7 Dynamics of aboveground phytomass of the circumpolar Arctic tundra during the past three  
8 decades, *Environ. Res. Lett.*, 7, 015506, 2012.
- 9 Esau, I., Davy, R., and Outten, S.: Complementary explanation of temperature response in the  
10 lower atmosphere, *Environ. Res. Lett.*, 7, 044026, 2012.
- 11 Frey, K., and Smith, L.: How well do we know northern land cover: Comparison of four global  
12 vegetation and wetland products with a new ground-truth database for West Siberia, *Global*  
13 *Biogeochemical Cycles*, 21, GB1016, 2007.
- 14 [Frost, G., Epstein, H., Walker, D., Matyshak, G., and Ermokhina, K.: Patterned-ground](#)  
15 [facilitates shrub expansion in Low Arctic tundra, \*Environmental Research Letters\*, 8, 015035,](#)  
16 [2013.](#)
- 17 Frost, G. V., and Epstein, H. E.: Tall shrub and tree expansion in Siberian tundra ecotones since  
18 the 1960s, *Global Change Biology*, 20, 1264–1277, 2014.
- 19 Goetz, S. J., Bunn, A. G., Fiske, G. J., and Houghton, R. A.: Satellite-observed photosynthetic  
20 trends across boreal North America associated with climate and fire disturbance, *Proc. Nat.*  
21 *Acad. Sci.*, 102, 13,521–13,525, 2005.
- 22 Guay, K., Beck, P., Berner, L., Goetz, S., Baccini, A., and Buermann, W.: Vegetation  
23 productivity patterns at high northern latitudes: a multi-sensor satellite data assessment, *Global*  
24 *Change Biology*, 20, 3147–3158, 2014.
- 25 Gutman, G., and Groisman, P. (eds.): *Environmental changes in Siberia: Regional changes and*  
26 *their consequences*, Springer, Amsterdam, 2013.
- 27 Hicks Pries, C. E. H., Schuur, E. A. G., Vogel, J. G., and Natali, S. M.: Moisture drives surface  
28 decomposition in thawing tundra, *Journal of Geophysical Research Biogeosciences*, 118, 1133-  
29 1143, 2013

- 1 Hinzman, L. D., Bettez, N. D., Bolton, W. R., et al. (32 co-authors): Evidence and implications  
2 of recent climate change in northern Alaska and other Arctic regions, *Climatic Change*, 72,  
3 251–298, 2005.
- 4 Ippolitov, I. I., Loginov, S. V., Kharyutkina, E. V., and Moraru, E. I.: Climate variability over  
5 the Asian territory of Russia during 1975–2012, *Geography and Natural Resources*, 35, 310–  
6 318, 2014.
- 7 Jeong, J.-H., Kug, J.-S., Kim, B.-M., Min, S.-K., Linderholm, H. W., Ho, C.-H., Rayner, D.,  
8 Chen, D., and Jun, S.-Y.: Greening in the circumpolar high-latitude may amplify warming in  
9 the growing season, *Clim. Dyn.*, 38, 1421–1431, 2011.
- 10 [Juszk, I., Eugster, W., Heijmans, M., and Schaepman-Strub, G.: Contrasting radiation and soil](#)  
11 [heat fluxes in Arctic shrub and wet sedge tundra, \*Biogeosciences Discussions\*, 2016, doi:](#)  
12 [10.5194/bg-2016-41](#)
- 13 Keeling, C. D., Chin, J. F. S., and Whorf, T. P.: Increased activity of northern vegetation  
14 inferred from atmospheric CO<sub>2</sub> measurements, *Nature*, 382, 146–149, 1996.
- 15 Kornienko, S., and Yakubson, K.: A study of transformation of vegetation cover in over some  
16 areas of Taz peninsula using satellite imagery, *Arctic: Ecology and Economics*, 4, 46–57, 2011.
- 17 Koronatova, N., and Milyaeva, E.: Plant community succession in post-mined quarries in the  
18 northern-taiga zone of West Siberia, *Contemporary Problems of Ecology*, 4(5), 513–518, 2011.
- 19 Kumpula, T., Pajunen, A., Kaarlejarvi, E., Forbes, B. C., and Stammer, F.: Land use and land  
20 cover change in Arctic Russia: Ecological and social implications of industrial development,  
21 *Glob. Environ. Change*, 21(2), 550–562, 2011.
- 22 Kumpula, T., Forbes, F., Stammer, F., and Meschyb, N.: Dynamics of a coupled system:  
23 Multi-resolution remote sensing in assessing social-ecological responses during 25 years of gas  
24 field development in Arctic Russia, *Remote Sensing*, 4, 1046–1068, 2012.
- 25 Lapenis, A., Shvidenko, A., Shepaschenko, D., Nilsson, S., and Aiyyer, A.: Acclimation of  
26 Russian forests to recent changes in climate, *Global Change Biology*, 11, 2090–2102, 2005.
- 27 Leibman, M., Khomutov, A., Gubarkov, A., Mullanurov, D., and Dvornikov, Y.: The research  
28 station “Vaskiny Dachi”, Central Yamal, West Siberia, Russia – A review of 25 years of  
29 permafrost studies, *Fennia*, 193, 3–30, 2015.

- 1 [Lloyd, A., Yoshikawa, K., Fastie, C.L., Hinzman, L., and Fraver, M.: Effects of permafrost](#)  
2 [degradation on woody vegetation at arctic treeline on the Seward Peninsula, Alaska, \*Permafrost\*](#)  
3 [and \*Periglacial Processes\*, 14\(2\), 93-101, 2003.](#)
- 4 Lloyd, A., and Bunn, A.: Responses of the circumpolar boreal forest to 20th century climate  
5 variability, *Environ. Res. Lett.*, 2, 045013, 2007.
- 6 MacDonald, G., Kremenetski, K., and Beilman, D.: Climate change and the northern Russian  
7 treeline zone, *Phil. Trans. Roy. Soc. B*, 363, 2283–2299, 2008.
- 8 Macias-Fauria, M., Forbes, B., Zetterberg, P., and Kumpula, T.: Eurasian Arctic greening  
9 reveals teleconnections, *Nature Climate Change*, 2(8), 613–618, 2012.
- 10 Melnikov, E. S., Leibman, M. O., Moskalenko, N. G., and Vasiliev, A. A.: Active-layer  
11 monitoring in the cryolithozone of West Siberia, *Polar Geography*, 28, 267–285, 2004.
- 12 Moskalenko, N.: Impact of climate warming on vegetation cover and permafrost in West  
13 Siberia northern taiga, *Natural Science*, 5(1A), 144–148, 2013.
- 14 Myneni, R. B., Keeling, C. D., Tucker, C. J., Asrar, G., and Nemani, R. R.: Increased plant  
15 growth in the northern high latitudes from 1981 to 1991, *Nature*, 386, 698–702, 1997.
- 16 Nauta, A. L., Heijmans, M., Blok, D., Limpens, J., Elberling, B., Gallagher, A., Li, B., Petrov,  
17 R. E., Maximov, T. C., van Huissteden, J., and Berendse, F.: Permafrost collapse after shrub  
18 removal shifts tundra ecosystem to a methane source, *Nature Climate Change*, 5, 67-70, 2015.
- 19 Outten, S., and Esau, I.: A link between Arctic sea ice and recent cooling trends over Eurasia,  
20 *Climatic Change*, 110, 1069–1075, 2011.
- 21 Outten, S., Davy, R., and Esau, I.: Eurasian winter cooling: Intercomparison of Reanalyses and  
22 CMIP5 data sets, *Atmos. Ocean. Sci. Lett.*, 6(5), 324–331, 2013.
- 23 Pavlov, A. V., and Moskalenko, N. G.: The thermal regime of soils in the north of Western  
24 Siberia, *Permafrost and Periglacial Processes*, 13, 43–51, 2002.
- 25 Reynolds, M. K., Walker, D. A., Epstein, H. E., Pinzon, J. E., and Tucker, C. J.: A new estimate  
26 of tundra-biome phytomass from trans-Arctic field data and AVHRR NDVI, *Remote Sens.*  
27 *Lett.*, 3, 403–411, 2012.
- 28 Reynolds, M. K., Walker, D. A., Ambrosius, K. J., Brown, J., Everett, K. R., Kanevskiy, M.,  
29 Kofinas, G. P., Romanovsky, V. E., Shur, Y., and Webber, P. J.: Cumulative geocological

1 effects of 62 years of infrastructure and climate change in ice-rich permafrost landscapes,  
2 Prudhoe Bay Oilfield, Alaska, *Global Change Biology*, 20(4), 1211–1224, 2014.

3 Shulgina, T., Genina, E., and Gordov, E.: Dynamics of climatic characteristics influencing  
4 vegetation in Siberia, *Environ. Res. Lett.*, 6, 045210, 2011.

5 Shuman, J. K., and Shugart, H. H.: Evaluating the sensitivity of Eurasian forest biomass to  
6 climate change using a dynamic vegetation model, *Environ. Res. Lett.*, 4, 045024, 2009.

7 Solano, R. et al.: MODIS Vegetation Index User's Guide (MOD13 Series), 2010, Available at:  
8 <http://vip.arizona.edu>.

9 Sorokina, N.: Anthropological changes in the northern taiga ecosystems of West Siberia  
10 (Nadym area), Ph.D. thesis, Institute for the Earth's Cryosphere, 2003

11 Streletskiy, D., Shiklomanov, N., and Nelson, F.: Permafrost, infrastructure, and climate  
12 change: a GIS-based landscape approach to geotechnical modeling, *Arctic, Antarctic and*  
13 *Alpine Res.*, 44(1), 368–380, 2012.

14 Sturm, M., McFadden, J., Liston, G., II, F. C., Racine, C., and Holmgren, J.: Snow–shrub  
15 Interactions in Arctic tundra: A hypothesis with climatic implications, *J. Clim.*, 14, 336–343,  
16 2001.

17 Urban, M., Forkel, M., Eberle, J., Hüttich, C., Schmullius, C., and Herold, M.: Pan-Arctic  
18 climate and land cover trends derived from multi-variate and multi-scale analyses (1981–2012),  
19 *Remote Sensing*, 6, 2296–2316, 2014.

20 Tang, Q., and Leng, G.: Damp summer warming accompanied with cloud cover increase over  
21 Eurasia from 1982 to 2009, *Environ. Res. Lett.*, 7, 014004, 2012.

22 Walker, D., Leibman, M., Epstein, H., et al. (17 co-authors): Spatial and temporal patterns of  
23 greenness on the Yamal Peninsula, Russia: interactions of ecological and social factors  
24 affecting the Arctic normalized difference vegetation index, *Environ. Res. Lett.*, 4, 045004,  
25 2009.

26 Walker, M., Wahren, C., Hollister, R., et al.: Plant community responses to experimental  
27 warming across the tundra biome. *Proceed. Nat. Acad. Sci.*, 5(103), 1342–1346, 2006.

28 Yakubson, K., Kornienko, S., Razumov, S., Dubrovin, V., Kritsuk, L., and Yastreba, N.:  
29 Geoindicators of the environmental change in areas of intensive oil and gas development and  
30 methods of their evaluation, *Georesources, geoenergetics, geopolitics*, 2, 2012.

- 1 Yu, Q., Epstein, H. E., Engstrom, R., Shiklomanov, N., and Streletskiy, D.: Land cover and  
2 land use changes in the oil and gas regions of Northwestern Siberia under changing climatic  
3 conditions, *Environmental Research Letters*, 10, 124020, 2015.
- 4 Zhang, X., Friedl, M. A., Schaaf, C. B., Strahler, A. H., and Schneider, A.: The footprint of  
5 urban climates on vegetation phenology, *Geophys. Res. Lett.*, 31, L12209, 2004.
- 6 Zhao, T., Bergen, K., Brown, D., and Shugart, H.: Scale dependence in quantification of land-  
7 cover and biomass change over Siberian boreal forest landscapes, *Landscape Ecology*, 24,  
8 1299–1313, 2009.
- 9 Zhou, L., Tucker, C., Kaufmann, R., and Slayback, D.: Variations in northern vegetation  
10 activity inferred from satellite data of vegetation index during 1981 to 1999, *J. Geophys. Res.*,  
11 106(D17), 20,069–20,083, 2001.

1 **Table 1.** List of 28 urban and industrial areas (cities) considered in this study. The city  
 2 population (pop. in thousands inhabitants) is estimated according to the Russian national census  
 3 for 2010. The mean background *NDVImax* (nature) is given for the most distant 40-km ring.  
 4 The relative trends are given for the time series without the years with the minimum and  
 5 maximum *NDVImax*. Statistically significant trends at 95% level are underlined. The biomes  
 6 are abbreviated as: tundra (T); forest-tundra (FT); norther taiga forest (NTF); and middle taiga  
 7 forest (MTF).

8

N	City Name	Coord.	Pop. (x1000)	Biome	Mean NDVImax (nature)	NDVImax trend (nature) [% dec <sup>-1</sup> ]	Mean NDVImax (city core)	NDVImax trend (city core) [% dec <sup>-1</sup> ]
1	Beloyarsky	63°42'N 66°40'E	↓49	NTF	0.78	-0.6%	0.65	+3.0%
2	Bovanenkovskiy	70°21'N 68°32'E	2 - 6	T	0.65	+1.4%	0.57	<u>-10%</u>
3	Gubkinsky	64°26'N 76°27'E	↓26	FT	0.65	-1.1%	0.53	+1.1%
4	Khanty-Mansyisk	61°48'N 69°10'E	↑93	MTF	0.78	-2.6%	0.62	-3.0%
5	Kogalym	62°14'N 74°32'E	↑61	MTF	0.69	-1.3%	0.65	-2.0%
6	Labytnangi	66°39'N 66°25'E	↑26	FT	0.75	-2.0%	0.60	<u>+6.0%</u>
7	Langepas	61°15'N 75°10'E	↑43	MTF	0.77	-1.9%	0.72	+1.4%
8	Megion	61°22'N 76°06'E	↓49	MTF	0.77	+0.8%	0.52	<u>+6.9%</u>
9	Muravlenko	63°47'N 74°31'E	↓33	NTF	0.65	-1.0%	0.47	-3.6%
10	Nadym	65°32'N 72°31'E	↓46	NTF	0.71	+1.2%	0.55	-1.8%
11	Nefteugansk	61°55'N 72°36'E	↓126	MTF	0.76	-2.1%	0.56	-3.0%
12	Nizshnavtovsk	60°56'N 76°33'E	↓266	MTF	0.78	-1.8%	0.60	<u>-4.7%</u>
13	Novyi Urengoy	66°05'N 76°04'E	↓116	NTF	0.63	<u>+4.0%</u>	0.54	+3.8%
14	Noyabrsk	63°11'N 75°27'E	↓107	MTF	0.69	No change	0.60	+0.6%
15	Nyagan	62°08'N 65°24'E	↑56	MTF	0.79	-0.7%	0.59	<u>+4.6%</u>
16	Pangody	65°51'N 74°31'E	↑11	FT	0.68	+3.6%	0.49	-2.1%
17	Pokachy	61°44'N 75°35'E	↑17	MTF	0.74	-1.5%	0.53	+1.6%
18	Purpe	64°29'N 76°42'E	↓10	FT	0.67	-0.4%	0.49	<u>-4.3%</u>
19	Pyt-Yakh	60°44'N 72°49'E	↑41	MTF	0.81	-2.0%	0.67	+1.1%

20	Raduzhnyi	62°06'N 77°28'E	↓43	MTF	0.72	-1.5%	0.65	-0.3%
21	Salekhard	66°31'N 66°36'E	↑48	FT	0.75	-0.9%	0.54	<b>+5.5%</b>
22	Sovetskiy	61°21'N 63°34'E	↑28	MTF	0.79	-0.6%	0.57	0.9%
23	Surgut	61°15'N 73°24'E	↑332	MTF	0.75	-2.4%	0.53	-0.2%
24	Tarko-Sale	64°55'N 77°47'E	↑21	FT	0.69	-0.6%	0.51	<b>-5.9%</b>
25	Tazovskiy	67°28'N 78°42'E	↓7	NTF	0.73	+1.5%	0.63	<b>+7.6%</b>
26	Uray	60°07'N 64°46'E	↑40	MTF	0.76	-0.2%	0.67	+0.5%
27	Urengoy	65°57'N 78°21'E	↓11	FT	0.68	-1.0%	0.55	-2.1%
28	Yar-Sale	66°52'N 70°49'E	↓7	T	0.66	+3.9%	0.52	-2.8%

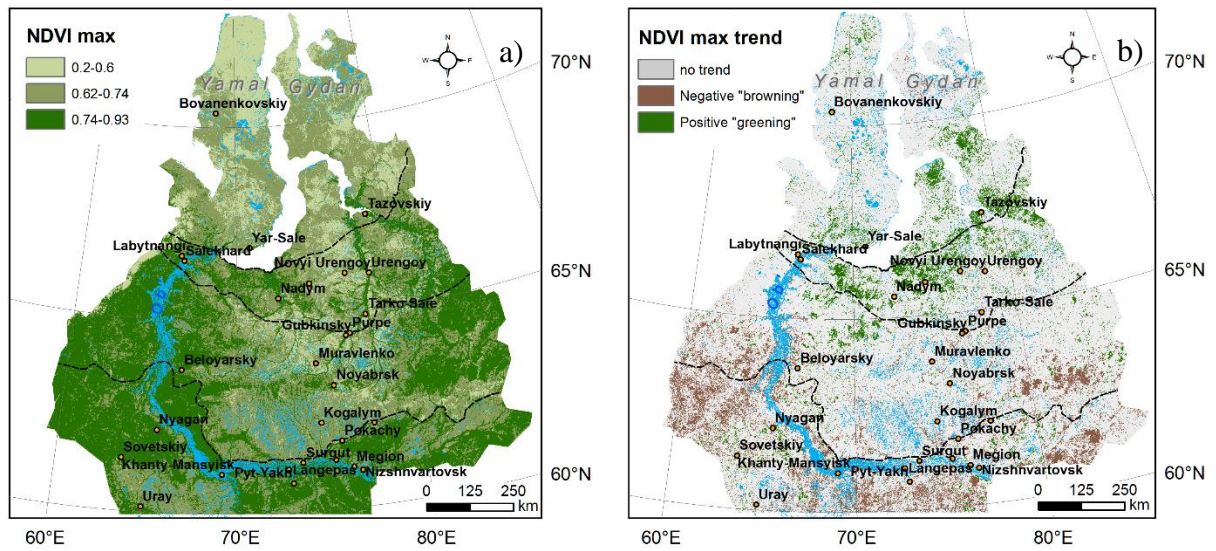
1

2 **Table 2.** Fractional areas of the *NDVImax* changes for 2000–2014 in four NWS biomes

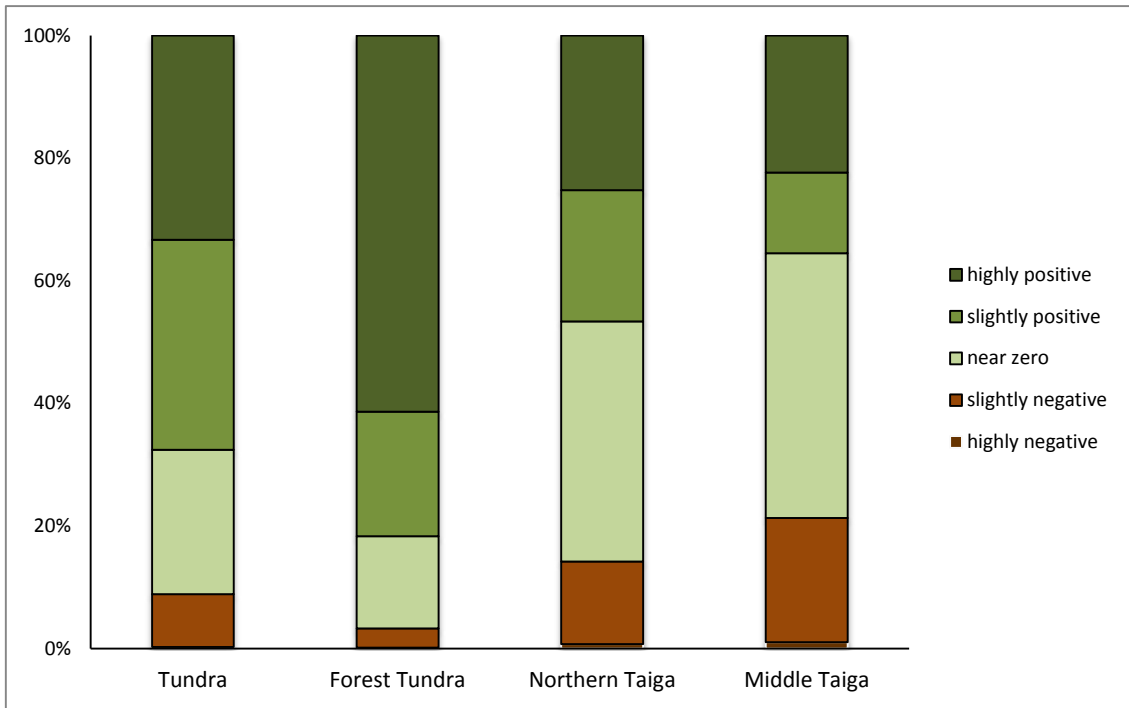
	<i>NDVImax</i> change	Tundra	Forest Tundra	Northern Taiga	Middle Taiga
Highly negative	≤ -0.006	0.28%	0.18%	0.73%	1.05%
Slightly negative	-0.006 to -0.003	8.58%	3.12%	13.48%	20.25%
Near zero	-0.003 to 0.003	23.56%	15.01%	39.19%	43.18%
Slightly positive	0.003 to 0.006	34.25%	20.33%	21.38%	13.16%
Highly positive	0.006 <	33.33%	61.36%	25.22%	22.36%

3



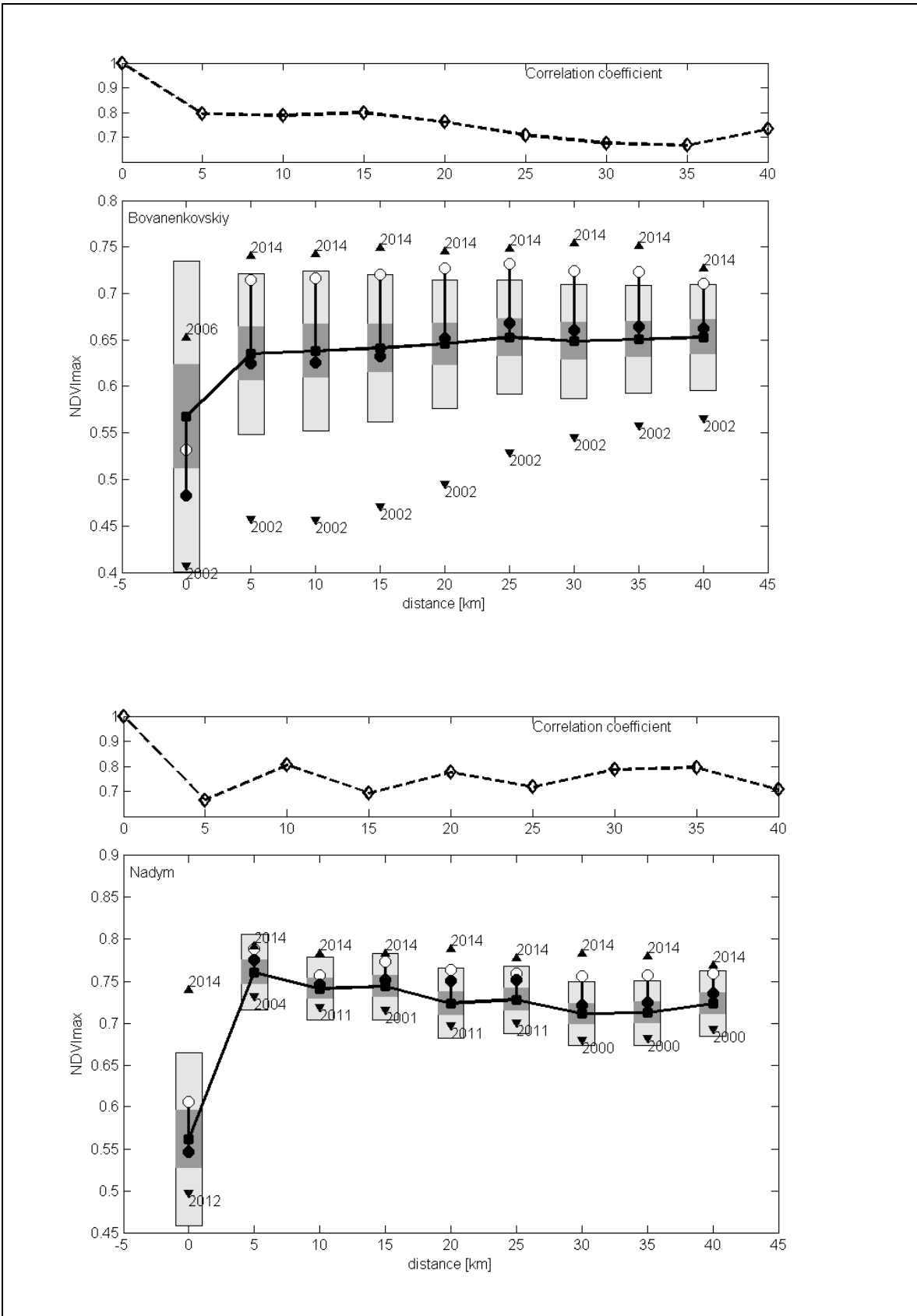


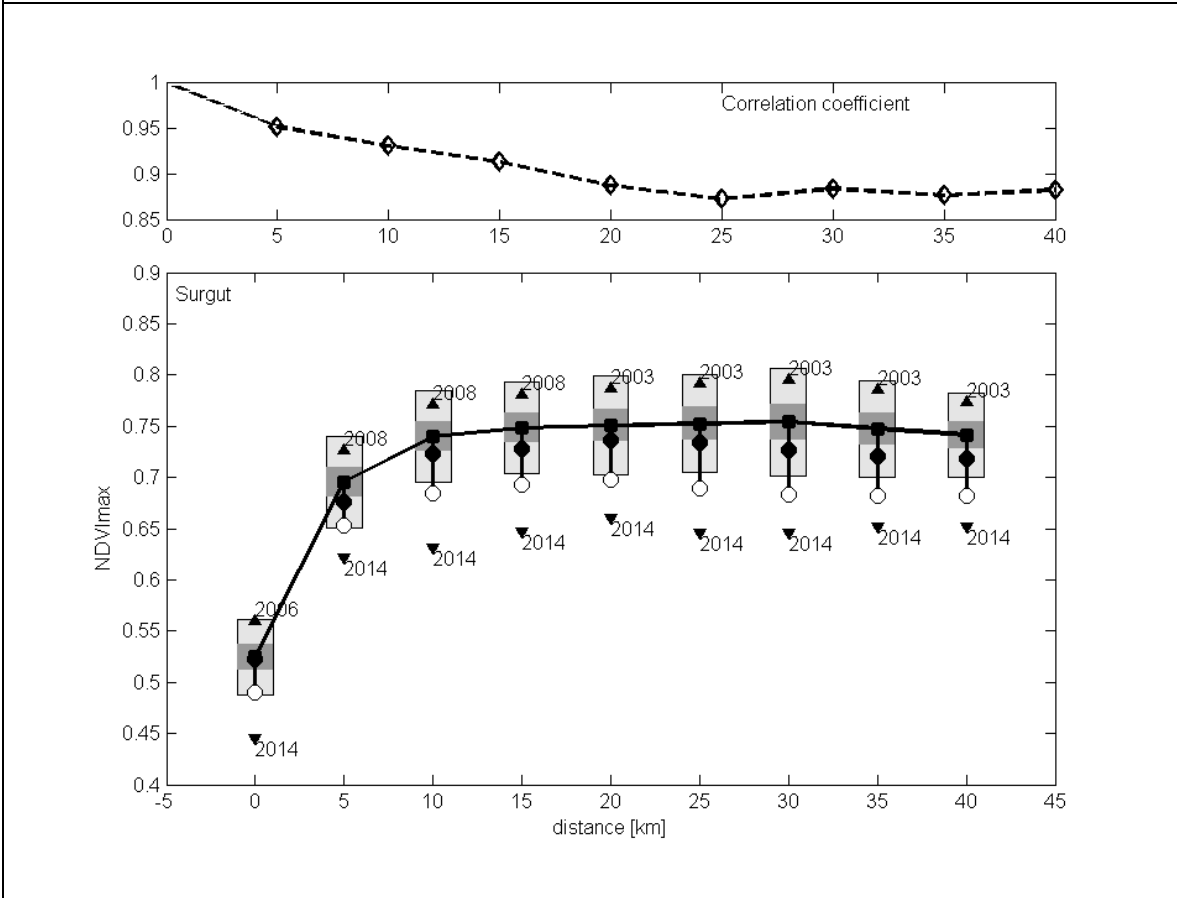
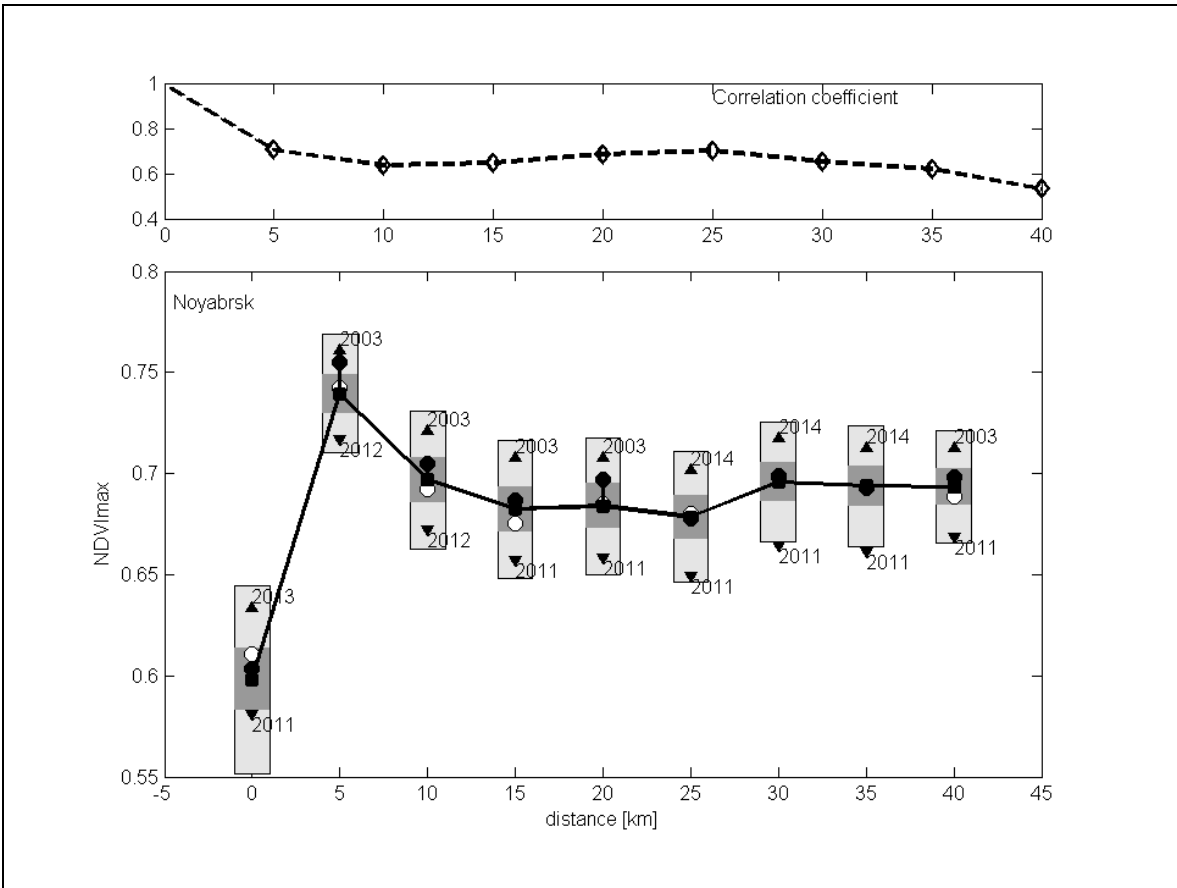
1 **Figure 1.** The 15-year mean absolute *NDVImax* (a) and the statistically significant (at  $\alpha < 0.05$ )  
 2 *NDVImax* trends for 2000-2014 (b).  
 3



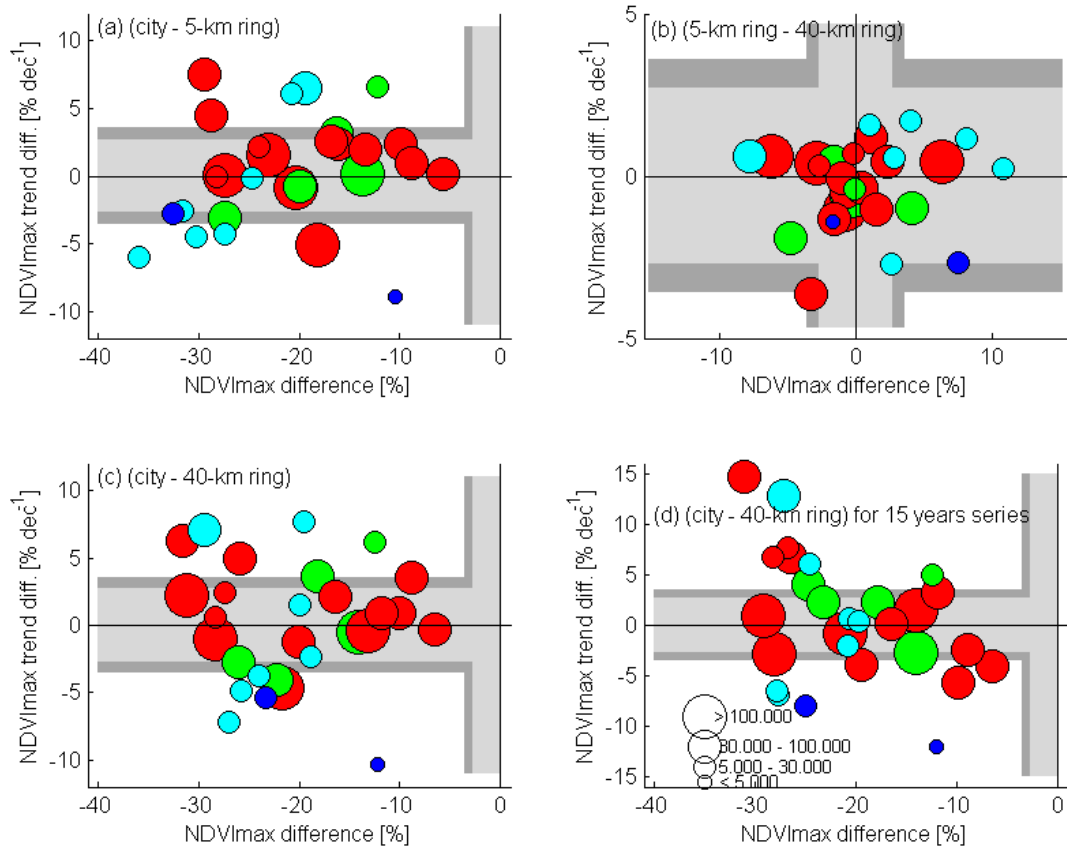
1 **Figure 2.** Total fractional areas of *NDVImax* changes for classes as defined in Table 2.

2





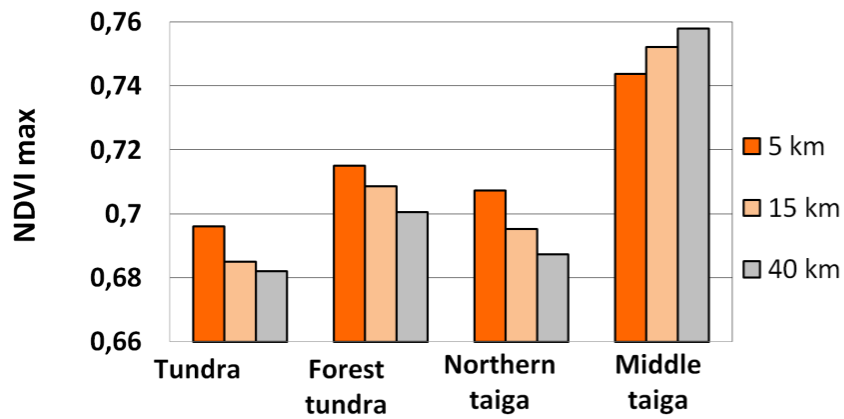
1 **Figure 3.** The statistical structure of *NDVImax*, correlations and trends at the city-core and the  
2 8 rings in the buffer zones of Bovanenkovskiy (tundra), Nadym (tundra-forest), Noyabrsk  
3 (northern taiga) and Surgut (middle taiga). The distance is given in km from the central pixel  
4 of each city core. The upper panels for each city show the decay of correlations (extreme years  
5 were excluded from the calculation) between the *NDVImax* variations at the city core and in the  
6 corresponding distance. The lower panels show: bold line with squares – the mean *NDVImax*;  
7 dark gray rectangle – one standard deviation of *NDVImax* for each ring; light gray rectangles –  
8 three standard deviations; triangles – the years with the maximum (upward-looking triangle)  
9 and the minimum (downward-looking triangle) of *NDVImax*; vertical black line with white  
10 circle – the magnitude of the *NDVImax* change obtained as the trend multiplied by 15 years;  
11 black circle – the same as the white circle but for the trends obtained when the maximum and  
12 the minimum *NDVImax* were excluded.  
13



1

2 **Figure 4.** The relative changes  $\Delta N(i, j)$  ( $x$ -axis) and  $\Delta R(i, j)$  ( $y$ -axis) given in % for (a) the city  
 3 core ( $i = 0$ ) and the closest 5-km ring ( $j = 1$ ); (b) the closest ( $i = 1$ ) and the most distant  
 4 background ( $j = 8$ ) rings; and (c) the city core ( $i = 0$ ) and the ring ( $j = 8$ ). The numbers were  
 5 obtained for the time series without the years of the maximum and minimum  $NDVImax$ . Panel  
 6 (d) shows the same as (c) but for the full 15-year time series. The circle color indicates: blue –  
 7 tundra; cyan – forest-tundra; green – northern taiga; red – middle taiga. The circle size indicates  
 8 the city population in 2005 as shown on panel (d). The gray shading shows the statistical  
 9 confidence envelopes (99% – dark gray shading; 95% – light gray shading).

10



1  
 2 **Figure 5.** Aggregated *NDVI*<sub>max</sub> for 5-km (closest to the city core), 15-km and 40-km (the most  
 3 distant background) rings around the cities in the corresponding four biomes.



5  
 6 **Figure 6.** Illustration of the alternative ecosystem development on anthropogenically disturbed  
 7 patches. The vegetation cover changes along the gas pipeline Nadym-Punga.

8