



1 **Direct observations of organic aerosols in common**  
2 **wintertime hazes in North China: insights into their size,**  
3 **shape, mixing state, and source**

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19



## 20 Abstract

21 Many studies have focused on the physicochemical properties of aerosol particles in  
22 unusually severe haze episodes instead of the more frequent and less severe hazes.  
23 Consistent with this lack of attention, the morphology and mixing state of organic  
24 matter (OM) particles in the frequent light and moderate (L&M) hazes in winter in  
25 North China Plain (NCP) have not been examined, even though OM dominates these  
26 fine particles. In the present work, morphology, mixing state, and size of organic  
27 aerosols in the L&M hazes were systematically characterized using transmission  
28 electron microscopy coupled with energy-dispersive X-ray spectroscopy, atomic force  
29 microscopy, and nanoscale secondary ion mass spectrometer, with the comparisons  
30 among an urban site (Jinan, S1), a mountain site (Tai, S2), and a background island  
31 site (Changdao, S3) in the same hazes. Based on their morphology, the OM particles  
32 were divided into six different types: spherical (type 1), near-spherical (type 2),  
33 irregular (type 3), domelike (type 4), dispersed-OM (type 5), and OM-coating (type 6).  
34 In the three sampling sites, type 1-3 of OM particles were most abundant in the L&M  
35 hazes and most of them were internally mixed with non-OM particles. The abundant  
36 near-spherical OM particles with higher sphericity and lower aspect ratio indicate that  
37 these primary OM particles formed in cooling, polluted plumes from coal combustion  
38 and biomass burning. Based on the Si-O-C ratio in OM particles, we estimated that 71%  
39 of type 1-3 OM particles were associated with coal combustion. Our result suggests  
40 that coal combustion in residential stoves was a widespread source from urban to rural  
41 areas in the NCP. Average OM thickness which correlates with the age of the air  
42 masses in type 6 particles only slightly increased from S3 to S2 to S1, suggesting that  
43 the L&M hazes were usually dry (relative humidity < 60%) with weak  
44 photochemistry and heterogeneous reactions between aerosols and gases. We  
45 conclude that the direct emissions from these coal stoves without any pollution  
46 controls in rural areas and in urban outskirts mainly contribute into the regional L&M  
47 hazes in North China.

48



## 49 1 Introduction

50 Atmospheric particulate matter is composed of diverse chemical compounds,  
51 both organic and inorganic matters. Organic aerosol particles are of two types:  
52 primary organic aerosol (POA), directly emitted from fossil fuel combustion, biomass  
53 burning, vehicular exhaust, and cooking; and secondary organic aerosol (SOA),  
54 formed from the oxidation of gaseous volatile organic compounds (Kanakidou et al.,  
55 2005). Organic aerosols account for 18-70% of the non-refractory submicron aerosol  
56 particles in the atmosphere (Zhang et al., 2007). It is well known that organic aerosols  
57 affect the atmosphere through the interaction with reactive trace gases, water vapor,  
58 clouds, precipitation, and radiation (Fuzzi et al., 2006). Organic aerosols also  
59 influence the physical and chemical properties of atmospheric particles (e.g., size,  
60 light-absorptivity, and hygroscopicity); they directly affect visibility and climate by  
61 scattering and absorbing solar radiation (Poschl, 2005; Kanakidou et al., 2005; Kulmala  
62 et al., 2004). Although most organic aerosol components are known to have a cooling  
63 effect on global climate, brown carbon in organic aerosols can absorb solar radiation  
64 at shorter wavelengths and lead to warming (Alexander et al., 2008). Moreover, many  
65 organic compounds (e.g., benzene, polycyclic aromatic hydrocarbons (PAHs),  
66 toluene), which are toxic to human and other biological species have been found in  
67 atmospheric particles (Mauderly and Chow, 2008).

68 In recent years, haze episodes have become one of the most serious environment  
69 problems in China, following the rapid urbanization and population growth in eastern  
70 China. China's Ministry of Environmental Protection on 1 January 2013, started to  
71 monitor daily  $PM_{2.5}$  concentrations and defined the various air pollution levels as:  
72 excellent ( $0\sim 35 \mu\text{g m}^{-3}$ ), good ( $35\sim 75 \mu\text{g m}^{-3}$ ), light ( $75\sim 115 \mu\text{g m}^{-3}$ ), moderate  
73 ( $115\sim 150 \mu\text{g m}^{-3}$ ), heavy ( $150\sim 250 \mu\text{g m}^{-3}$ ), and severe ( $> 250 \mu\text{g m}^{-3}$ ) (HJ 663-2012).  
74 Haze as a weather phenomenon is defined by visibility  $\leq 10$  km and RH  $\leq 95\%$ .  
75 Previous studies have shown that haze levels normally are associated with  $PM_{2.5}$   
76 concentrations and RH (Shen et al., 2015; Wang et al., 2006; Chen et al., 2014). Based  
77 on their results, we classify severe haze days ( $< 5$  km) with  $PM_{2.5}$  concentrations  $\geq$



78  $250 \mu\text{g m}^{-3}$  and light (8-10 km) to moderate (5-8 km) haze days at  $75\text{-}250 \mu\text{g m}^{-3}$ , both  
79 with  $\text{RH} < 80\%$ .

80 Severe wintertime haze pollution episodes were driven to a large extent by  
81 secondary aerosol formation, and SOA were found to be of similar importance with  
82 secondary inorganic aerosols (Guo et al., 2014;Huang et al., 2014). However, Sun et  
83 al. (2013) showed that OM was dominant in wintertime haze pollution episodes but  
84 POA concentrations were much higher than SOA in the aerosol particles. The reason  
85 about the different results of OM particles among these studies could be the difference  
86 of air pollution levels and RH during the haze pollution episodes. Severe winter hazes  
87 were driven by stable synoptic meteorological conditions with high relative humidity  
88 (i.e.,  $\text{RH} > 80\%$  frequently or consistently occurred during the haze episodes) (Tie et  
89 al., 2015). Due to the dimming effect of high-concentration aerosol particles, an  
90 enhanced production rate of secondary aerosols was suggested as an important  
91 contribution from the complicated heterogeneous reactions (Li et al., 2011c;Zheng et  
92 al., 2015). Because of severe haze episodes with unusually high  $\text{PM}_{2.5}$  concentrations  
93 ( $> 250 \mu\text{g m}^{-3}$ ) and low visibility ( $< 5 \text{ km}$ ), many scientists have been investigated the  
94 physicochemical properties of their aerosol particles (Huang et al., 2014;Guo et al.,  
95 2014;Zheng et al., 2015). Although this knowledge is critical to understand severe  
96 haze formation and its impacts on human health, the frequency of severe haze events  
97 is low, they are of short duration. For example, we statistically analyzed haze days  
98 during the winter ( $\sim 92$  days) of 2014-2015 in nine different cities. Figure S1 shows  
99 that light and moderate (L&M) haze days occurred 22-63% of the time and that severe  
100 haze days were less frequent at 4-32%, with the variation dependent on location  
101 within the NCP. As we know, the L&M hazes are precursors of severe ones. Zheng et  
102 al. (2015) suggested that characteristics of aerosol particles in severe hazes would not  
103 be the same in L&M hazes. Compared to severe hazes, the L&M hazes were most  
104 frequent in winter and lasted longer in the NCP. Therefore, understanding aerosol  
105 particles in the more common L&M hazes in the NCP is important to further evaluate  
106 their impacts on human health and regional climate.



107 Various “bulk” analytical instruments have been used to study organic aerosol  
108 particles during haze episodes. High resolution time-of-flight aerosol mass  
109 spectrometry (HR-AMS) was applied to determine the mass concentrations and bulk  
110 composition of organic aerosols (Sun et al., 2010). Gas chromatography-mass  
111 spectrometry (GC-MS) provided chemical composition and structures of organics in  
112 aerosols (Fu et al., 2012; Wang et al., 2009). It should be noted that bulk analytical  
113 techniques only provide average properties of PM<sub>2.5</sub> and the mixing state, phase, and  
114 morphology of organic particles remain unknown. Detailed information about  
115 individual organic particles, moreover, is critical to evaluate their formation, their  
116 sources, and their hygroscopic and optical properties in the atmosphere. For example,  
117 copious tar balls containing homogeneous BrC occur in the smoldering smoke from  
118 biofuels (Chakrabarty et al., 2010; Adachi and Buseck, 2011; Alexander et al.,  
119 2008; Chakrabarty et al., 2013; China et al., 2013; Hand et al., 2005; Posfai et al., 2004).  
120 Atmospheric particles undergo liquid-liquid phase separations and go on to form OM  
121 coatings on inorganic aerosol particles (You et al., 2012). The surface coating by OM  
122 on individual particles influences water uptake and evaporation of individual  
123 particles and their heterogeneous reactions in the atmosphere (Shiraiwa et al.,  
124 2011; Zawadowicz et al., 2015; Riipinen et al., 2011). Despite the importance of these  
125 phenomena, the morphology and mixing state of OM particles in wintertime L&M  
126 hazes in the NCP have not been examined, although OM is dominant in fine particles  
127 (Sun et al., 2013).

128 To characterize organic aerosols in greater detail in L&M hazes, individual  
129 particles in the NCP in winter were analyzed using transmission electron microscopy  
130 coupled with energy-dispersive X-ray spectroscopy (TEM/EDX), atomic force  
131 microscopy (AFM), and Nanoscale secondary ion mass spectrometer (NanoSIMS).  
132 Aerosol particles were simultaneously collected on TEM grids in the NCP during  
133 13-23 December, 2014. Morphology, mixing state, and size of organic aerosols were  
134 systematically characterized and compared at the three sampling sites (background  
135 island site, mountain site, and urban site) in the same haze. This information enables



136 the discussion of source and ageing mechanisms of OM particles, which leads to  
137 insights about the formation of regional wintertime L&M hazes in the NCP.

## 138 2 Experimental Methods

### 139 2.1 Sampling sites and particle collection

140  $PM_{2.5}$  and individual particle samples were simultaneously collected during  
141 13-23 December, 2014 at three sampling sites: an urban site (S1), a mountain site (S2),  
142 and a background site (S3) in the NCP (Fig. 1). S1- the urban site in Jinan (36.67 °N,  
143 116.98 °E) is 50 km north of S2. S1 is a typical polluted city with high-density  
144 residential areas surrounded by large industrial zones (Li et al., 2011c). Aerosol  
145 particles collected at S1 mainly reflect local, ground-based urban and industrial  
146 emissions. S2-Mt. Tai (1534 m a.s.l., 36.251 °N, 117.101 °E) is the highest mountain in  
147 the middle of the NCP, ~230 km inland from the Bohai and Yellow Seas. S2 is the  
148 perfect location to observe air pollutants near the planetary atmospheric layer over the  
149 NCP. Aerosol particles collected at S2 represent regional transport (Li et al., 2011a).  
150 S3-Changdao Island, the National Station for Background Atmospheric Monitoring  
151 site (38.19 °N, 120.74 °E), is in the Bohai Sea. During the winter monsoon season, S3  
152 is the downwind of the Jing-jin-ji area (i.e., Beijing city, Tianjin city, and Hebei  
153 province) and Shandong province. Therefore, S3 serves as a polluted background site  
154 from the transport of continental air. Therefore, aerosol particles collected at the three  
155 sampling sites display the different pollutant characteristics of polluted urban air,  
156 upper air layer, and background island air in the NCP.

157  $PM_{2.5}$  was collected on 90 mm quartz filters for 9-11.5 h using three KB-120  
158 samplers at a flow rate of 100 L/min. The quartz filters were stored in a refrigerator  
159 for OC, EC, and water soluble ion analysis. In the study, OC and EC concentrations of  
160 70 quartz filters were analyzed by an OC/EC analyzer (Sunset Lab) and water-soluble  
161 ions (i.e.,  $K^+$ ,  $Na^+$ ,  $Ca^{2+}$ ,  $Mg^{2+}$ ,  $NH_4^+$ ,  $F^-$ ,  $SO_4^{2-}$ ,  $NO_3^-$ , and  $Cl^-$ ) by an ion  
162 chromatography system (Dionex ICs-90). Three single-stage cascade impactors with a  
163 0.5-mm diameter jet nozzle at a flow rate of 1.0 L/min were used to collect aerosols



164 onto copper TEM grids coated with carbon film (carbon type-B, 300-mesh copper,  
165 Tianld Co., China). The collection efficiency of the impactor is 50% for particles with  
166 an aerodynamic diameter of 0.25  $\mu\text{m}$  and with a density of 2  $\text{g cm}^{-3}$  (Li et al., 2011a).  
167 After sample collection, the Cu grids were placed in a sealed, dry and clean  
168 environment until the TEM analysis. 11 aerosol samples at each sampling site were  
169 selected and analyzed by the TEM.

## 170 2.2 TEM analyses

171 The JEOL JEM-2100 transmission electron microscopy operated at 200 kV with  
172 energy-dispersive X-ray spectrometry (TEM/EDX) was used to analyze individual  
173 particles. An energy-dispersive X-ray spectrometer (EDX) can detect elements  
174 heavier than carbon. EDX spectra were executed for 15 s to minimize the potential  
175 beam damage. TEM grids are made of copper (Cu), so the Cu element will be  
176 excluded in the analyses. The distribution of particles on the TEM grids was not  
177 uniform: coarser particles were deposited near the center and finer particles dispersed  
178 on the fringe. To make sure that the analyzed particles were representative of the  
179 entire size range, three to four areas were chosen from the center and periphery of the  
180 sampling spot on each sample.

## 181 2.3 NanoSIMS analysis

182 Individual aerosol particles were analyzed using a nanoscale secondary ion mass  
183 spectrometer (NanoSIMS) 50L, an ultrahigh vacuum technique for surface and  
184 thin-film analysis at the Institute of Geology and Geophysics, Chinese Academy of  
185 Sciences. In this study,  $^{12}\text{C}^-$ ,  $^{16}\text{O}^-$ ,  $^{12}\text{C}^{14}\text{N}^-$ ,  $^{14}\text{N}^{16}\text{O}_2^-$ ,  $^{32}\text{S}^-$  ions in individual particles  
186 were obtained when the  $\text{Cs}^+$  primary ion beam caused the ionization of atoms within  
187 the particles. Furthermore, ion intensity mappings of individual particles with  
188 nanometer resolution can show the distribution of different ions.  $^{12}\text{C}^-$  and  $^{12}\text{C}^{14}\text{N}^-$   
189 represent the organic matter in individual particles (Chi et al., 2015; Ghosal et al.,  
190 2014; Li et al., 2016a).

## 191 2.4 AFM analysis



192 Atomic force microscopy (AFM) with a tapping mode analyzed aerosol particles  
193 under ambient conditions. AFM, a digital NanoscopeIIIa Instrument, can determine  
194 the three dimensional morphology of particles. The AFM settings contain imaging  
195 forces between 1 and 1.5 nN, scanning rates between 0.5 and 0.8 Hz, and scanning  
196 range sizes at 10  $\mu\text{m}$  with a resolution of 512 pixels per length. After the AFM  
197 analysis, composition of the same particles was confirmed by TEM, with 20, 25, and  
198 13 individual particles analyzed by this method for each of the three sampling sites.  
199 The Nanoscope analysis software can automatically obtain bearing area (A) and  
200 bearing volume (V) of each analyzed particle according to the formulas described by  
201 Chi (Chi et al., 2015).

202 The relationship is shown in Figure S2 (EVD=0.8334ECD (S1),  
203 EVD=0.7286ECD (S2), and EVD=0.6601ECD (S3)). Therefore, the equivalent circle  
204 diameter (ECD,  $x$ ) of individual aerosol particles measured from the iTEM software  
205 can be further converted into equivalent volume diameter (EVD,  $y$ ) based on these  
206 relationships.

207

### 208 3. Results

#### 209 3.1 Regional haze periods in North China Plain

210 Aerosol particles were collected in three regional L&M hazes during 13-23  
211 December, 2014 (Fig. S3). Moderate Resolution Imaging Spectroradiometer (MODIS)  
212 images clearly display a regional haze layer covering the three sampling sites in the  
213 NCP (Fig. 1). The average  $\text{PM}_{2.5}$  concentrations were  $96.6 \mu\text{g m}^{-3}$  (range:  $79\text{-}171 \mu\text{g m}^{-3}$ )  
214  $\text{m}^{-3}$ ),  $88.6 \mu\text{g m}^{-3}$  (range:  $76\text{-}110 \mu\text{g m}^{-3}$ ), and  $80.3 \mu\text{g m}^{-3}$  (range:  $75\text{-}84 \mu\text{g m}^{-3}$ ) on  
215 haze days, twice as high as on clear days ( $52$ ,  $48$ , and  $32.3 \mu\text{g m}^{-3}$ ) at S1, S2, and S3,  
216 respectively. The RH at all three sampling sites was lower than 60% during the  
217 sampling periods (Fig. S4).

218 Figure 1 shows similar back trajectories of air masses at S1, S2, and S3. Haze  
219 episodes started to form in the NCP when air masses changed from north to southwest  
220 or west. Therefore, the regional hazes in the NCP were controlled by the same



221 meteorological conditions (e.g., RH, temperature, and wind). The average  
222 concentrations of OC, EC, OC/EC, water-soluble ions and their mass proportions in  
223  $PM_{2.5}$  were much higher on haze days than on clear days at the three sampling sites  
224 (Table S1). OC on haze days was more than 1.7 times higher than that on clear days at  
225 three sites. OM concentration was estimated at 20-33  $\mu\text{g m}^{-3}$  and OM/ $PM_{2.5}$  ratio was  
226 at the range of 23-34% during haze days in the NCP (Table S1).

### 227 3.2 Morphology of organic particles

228 TEM/EDX identified five types of particles during the haze episodes: sulfates  
229 (including K-rich sulfate and ammonium sulfate), fly ash, mineral, soot, and  
230 C-dominated particles (Fig. S5). These results are consistent with previous studies  
231 during the haze episodes in the NCP (Li et al., 2012; Li et al., 2011c). In order to  
232 remove the interference of the carbon substrate on TEM grids, a nanoSIMS was  
233 employed to verify C-dominated particles through  $^{12}\text{C}^{14}\text{N}^-$  and  $^{12}\text{C}^-$  mappings (Fig. 2  
234 and Fig. S6). Figure 2 clearly shows that one near-spherical particle, which contains C,  
235 O, and minor Si on TEM grids, displays strong  $\text{CN}^-$  and  $\text{O}^-$  signals but no clear  $\text{NO}_2^-$   
236 and  $\text{S}^-$  signals. As a result, this type particle can be confirmed as the OM particle (Li  
237 et al., 2016a; Ghosal et al., 2014). TEM analysis showed that OM-containing particles  
238 were most abundant in all the haze samples, accounting for 70% of the 5090 analyzed  
239 particles (Fig. S7).

240 Based on the morphology of OM particles, they were divided into six different  
241 types: spherical (type 1, Fig. 3a), near-spherical (type 2, Fig. 3b), irregular (type 3,  
242 Fig. 3c), domelike (type 4, Fig. 3d), dispersed-OM (type 5, Fig. 3e), and OM-coating  
243 (type 6, Fig. 3f). Because the high-resolution TEM images of individual particles can  
244 clearly display particle interior mixing structures, it allows us to identify OM particles  
245 based on their different shapes in OM-containing particles (Fig. S5). Figure 4 shows  
246 that the proportions of type 1-3 in OM particles was 73%, following type 4 at 5%,  
247 type 5 at 13%, and type 6 at 14%. Further, we measured the projected area, the  
248 perimeter, the maximum projected length, and the maximum projected width of 967  
249 selected OM particles. From these data, the sphericity (Sph) and aspect ratio (AR) of



250 different types of OM particles were calculated, which characterize their shape and  
251 thereby imply their aging during transport and their emission sources (Li et al.,  
252 2016b). The Sph and AR were defined by the following formulas referred to by (Li et  
253 al., 2013).

254 *Aspect Ratio (AR)*. The maximum ratio of width and height of a bounding  
255 rectangle for the measured object is the aspect ratio. An aspect ratio of 1 (the lowest  
256 value) indicates a particle is not elongated in any direction.

257 *Sphericity (Sph)*. Sphericity describes the sphericity or “roundness” of the  
258 measured object by using central moments. A sphericity of 1 (the highest value)  
259 indicates a particle is perfectly spherical.

$$260 \quad Sph = \frac{\sqrt{4\pi S}}{P} = \frac{\sqrt{4\pi R_1^2}}{2\pi R_2} = \frac{R_1}{R_2} \quad (1)$$

$$261 \quad AR = \frac{L_{\max}}{W_{\max}} \quad (2)$$

262 Where S is projected area,  $R_1$  is equivalent area radius, P is perimeter,  $R_2$  is  
263 equivalent perimeter radius,  $L_{\max}$  is the maximum projected length, and  $W_{\max}$  is the  
264 maximum projected width.

265 Table 1 displays the Sph and AR of individual OM particles measured by the  
266 iTEM software. At the three different sampling sites, OM particles were in the fine  
267 range with diameters  $< 1\mu\text{m}$ . The statistics show that spherical OM particles  
268 exhibited the highest Sph at 0.96-0.99 and the lowest AR at 1.0-1.03 at the three sites,  
269 followed by OM coating (Sph: 0.88-0.93, AR: 1.06-1.08), near-spherical OM (Sph:  
270 0.82-0.83, AR: 1.12-1.13), domelike OM (Sph: 0.64-0.73, AR: 1.24-1.35), irregular  
271 OM (Sph: 0.50-0.57, AR: 1.39-1.48) and dispersed-OM particles (Sph 0.49-0.58, AR:  
272 1.35-1.46).

### 273 3.3 Mixing state of OM particles

274 Although we identified different types of OM particles in individual particles, 86%  
275 were internally mixed with non-OM particles, such as soot, mineral, metal, fly ash,  
276 and sulfate particles (Fig. S5). Based on their morphological mixing state, we



277 discriminated four OM internally mixed particles: OM-soot (Fig. 5a and Fig. S8),  
278 OM-mineral/metal (Fig. 5b), OM-fly ash (Fig. 5c), and OM-sulfate particles (Fig.  
279 5d-e). Our results show that 83% of type 1-4 OM particles were attached to soot,  
280 mineral, sulfate, and metal particles, only 17% of type 1-4 OM particles were  
281 externally mixed particles, and all the type 5-6 OM were internally mixed with sulfate  
282 particles (Fig. S7).

283 Figure 6 shows number fractions of OM internally mixed particles in different  
284 size bins from 0.04 to 4.5  $\mu\text{m}$  at the three sampling sites. OM-soot particles  
285 commonly occurred at S1 but they were mixed with certain amounts of sulfates at S2  
286 and S3 during the sampling period (Fig. S8). The major OM internally mixed particles  
287 include 45% OM-soot particles and 46% OM-sulfate particles at S1, 35% and 62% at  
288 S2, 33% and 56% at S3 (Fig. S7). As a result, the number fraction of OM-sulfate  
289 particles increases from S1 to S2 and S3. OM-soot containing particles dominated in  
290 the finer size range ( $< 300$  nm) at the three sampling sites (Fig. 6). In addition, 19% of  
291 OM-sulfate particles were internally mixed with inclusions (i.e., fly ash and metal) at  
292 all the three sampling sites (Fig. S7).

### 293 **3.4 Size distribution of OM-containing particles**

294 Figure 7 shows size distributions of OM-containing particles at the three  
295 sampling sites. Aerosol particles collected at S2 and S3 display a similar peak at  $\sim 400$   
296 nm, much smaller than the peak at 600 nm at the S1 site (Fig. 7a). This result  
297 indicates that sizes of locally emitted OM-containing particles are much larger than  
298 the long-range transported OM-containing particles. We further obtain size  
299 distributions of type 1-3 OM particles at the three sampling sites during one haze  
300 episode. Interestingly, type 1-3 OM particles displayed similar peaks around 350 nm  
301 at all three sampling sites (Fig. 7b). This result suggests that the OM sources were  
302 similar under different transport ranges in the same haze layer over the NCP.

303

## 304 **4. Discussion**

### 305 **4.1 Sources of OM-containing particles**



306 TEM adequately characterized the morphology and mixing state of  
307 OM-containing particles in wintertime L&M hazes. We found that the type 1-3 OM  
308 particles (Fig. 3a-c) were most abundant in the hazes and that most of them were  
309 internally mixed with non-OM particles. This result is consistent with one previous  
310 study which found abundant amorphous spherical OM particles in the outflow of a  
311 haze plume in East Asia (Zhu et al., 2013). Moreover, Li et al. (2012) found large  
312 amounts of type 1 OM particles in a coal-burning region in the China Loess Plateau in  
313 winter. However, some studies found abundant type 5-6 OM particles in the  
314 atmosphere. For example, Moffet et al. (2013) suggested that OM coating particles  
315 become dominant following particle transport from an urban to background site and  
316 they didn't report abundant type 1-3 OM particles in North America. Also, Adachi et  
317 al. (2014) only reported that most sulfate particles were coated by secondary OM  
318 coating in remote mountain air in Japan. Based on these comparisons, we conclude  
319 that those type 1-3 OM particles were not emitted by vehicular emissions in the NCP.

320 It should be noted that recent studies did not find abundant type 1-3 OM particles  
321 at three sampling sites on haze episodes in spring and summer (Li et al., 2011b; Yuan  
322 et al., 2015). Zhang et al. (2008) suggested that industrial boilers had cleaner  
323 combustion with much less by-product of particulate carbon and with much lower  
324 levels of OM, while residential stoves had significantly higher emissions of  
325 carbonaceous particulate matter with emission rates 100 times higher than that of  
326 industrial boilers. As a result, we believe that the type 1-3 OM particles were not  
327 emitted from heavy industries or coal-fired power plants but, instead, they were from  
328 coal combustion or biomass burning for household heating and cooking in wintertime.  
329 In particular, the abundant near-spherical OM particles with higher Sph and lower AR  
330 indicate that these OM particles formed in cooling, polluted plumes from coal  
331 combustion and biomass burning.

332 Biomass burning and coal combustion both can produce types 1-2 OM particles  
333 and all contain a certain amount of Si beside C and O (Li et al., 2012; Posfai et al.,  
334 2004; Hand et al., 2005; Adachi and Buseck, 2011). Li et al. (2012) found that primary



335 OM particles contain much higher Si from coal combustion than biomass burning. To  
336 evaluate OM sources in this study, we compared ratios of Si, O, and C in individual  
337 OM particles collected in haze and fresh OM particles from corn stalks and coal  
338 combustion conducted in the laboratory (supplementary materials). Figure 8 clearly  
339 shows that the Si ratio in individual OM particles is ordered as coal combustion >  
340 haze particles > corn stalks; and that 71% of haze OM particles are associated with  
341 coal combustion. Therefore, we can conclude that coal combustion contributes more  
342 type 1-3 OM particles than biomass burning in the wintertime L&M haze. This result  
343 is consistent with the source apportionment of OM particles based on their mass  
344 concentrations (Elser et al., 2016; Sun et al., 2013). Furthermore, the consistent Sph  
345 and AR of abundant type 1-3 OM particles at the three sampling sites (Table 1)  
346 suggests that coal combustion in residential stoves was a widespread emission source  
347 from urban to rural areas in the NCP. The vehicular emissions at S1 had a higher  
348 contribution of fine soot (i.e., BC) particles with diameters < 300 nm (Fig. 6). Fly  
349 ash/metal particles were normally considered from coal-fired power plant and heavy  
350 industries (Li and Shao, 2009; Shi et al., 2003; S. X. Wang, 2010). Although fly  
351 ash/metal-bearing particles occurred at three sampling sites but their fraction was only  
352 19% (Fig. S7), suggesting that these large stationary sources were not major sources  
353 for the primary particles.

#### 354 **4.2 Ageing of OM particles**

355 The complicated mixing structures of individual particles can be used to evaluate  
356 particle ageing mechanisms (Li et al., 2016b). Figure 1 suggests that S2 and S3 as the  
357 polluted background sites received aged particles after long-range transport and that  
358 the urban site of S1 received more fresh particles. Indeed, OM-soot particles at S2 and  
359 S3 sites were internally mixed with sulfates but not at S1 (Fig. S8). To evaluate  
360 particle ageing processes, we measured OM coating thickness in type 6 OM-coating  
361 particles (e.g., Fig. 5e) because coating thickness on secondary particles can be used  
362 to infer particle ageing during transport (Moffet et al., 2010; Moffet et al., 2013). In  
363 this study, OM coating thickness increased with particle size and that their average



364 values at the three sampling sites were ordered as  $S3 > S2 > S1$  (Fig. 9). The results  
365 suggest that coarse secondary particles underwent more ageing than the fine particles  
366 and that the particles at S3 underwent the most ageing. However, number fractions of  
367 type 6 OM particles were small at three sampling sites (14.8% at S2 and 12% at S3,  
368 and 2.9% at S1) (Fig. 9). The result indicates weak atmospheric reactions for SOA  
369 formation in the whole haze layer.

## 370 5. Conclusions and atmospheric implications

371 Abundant type 1-3 OM particles at S1, S2, and S3 suggested that coal  
372 combustion and biomass burning used for cooking and heating in residential sector  
373 in winter significantly contributed to the haze layer over the NCP. Although heavy  
374 industrial and coal-fired plants emitted large amount of gases such as  $SO_2$ ,  $NO_x$ , and  
375 VOCs in the NCP (Wang et al., 2012; Zhang et al., 2015; Wang et al., 2013a), we  
376 didn't observe many secondary organic and inorganic aerosols in wintertime L&M  
377 hazes; these aerosols are common particle types in heavy haze and fog episodes (Li et  
378 al., 2011c). The results indicated very weak photochemical reactions due to lower  $O_3$   
379 concentrations and weaker solar radiation, two constraints that reduce the conversion  
380 of the acidic gases into aerosol particles (Ma et al., 2012). Also, the L&M hazes were  
381 dry with  $RH < 60\%$  (Fig. S4), which prohibits heterogeneous reactions between  
382 aerosol and gases (Zheng et al., 2015). These reasons can explain why we found  
383 higher number and mass fractions of abundant type 1-3 primary OM particles and  
384 lower type 6 secondary OM particles in the hazes. This result is consistent with a  
385 previous study using AMS (Sun et al., 2013), which showed 69% primary OM  
386 particles and only 31% SOA in winter hazes. Concerning mixing processes, the sizes  
387 (Fig. 7) and shapes (Sph and AR in Table 1) of the type 1-3 OM particles were  
388 consistent with fresh particles at S1 and aged particles at S2 and S3. Therefore,  
389 coagulation between OM and non-OM particles was expected to be the major mixing  
390 mechanism on the dry haze with its weak photochemistry.

391 Once humidity exceeds 60% during a haze episode, secondary inorganic and  
392 organic species from heterogeneous reactions can significantly contribute to haze



393 aerosols (Zheng et al., 2015). In a word, primary emissions from cooking and heating  
394 likely caused the regional L&M haze episodes in the NCP. Moreover, coal  
395 combustion and biomass burning can emit K-rich sulfates (e.g., Fig. 5d) (Wang et al.,  
396 2013b). These residential coal stoves in rural areas and in the urban outskirts have no  
397 pollution controls and directly emit particulate carbon and other pollutants. The  
398 emission control of residential coarse coal combustion is simply not regulated by the  
399 national environmental protection bureau, even though this bureau has made much  
400 recent progress in controlling emissions from heavy industries and coal-fired power  
401 plants.

402 We also noticed that a large proportion of type 1-3 primary OM particles has been  
403 determined to be brown carbon (BrC), which can absorb shorter-wavelength solar  
404 radiation (Alexander et al., 2008). The BrC and BC in the haze layers over the NCP  
405 undoubtedly absorb solar energy and influence the vertical mixing of air. One  
406 modeling study shows that these absorbing aerosols can lead to a more stable  
407 atmospheric stratification over the NCP that decreases turbulence diffusion by 52%  
408 and decreases the planetary boundary layer (PBL) height by 33 % (Wang et al., 2015).  
409 OM particles from the direct emissions of coarse coal combustion and biomass  
410 burning not only contributed to PM<sub>2.5</sub> in the L&M hazes, but also influenced haze  
411 stability. Moreover, the type 1-3 OM particles from coal combustion in residential  
412 stoves mainly consist of PAHs (Zhang et al., 2008). In this study, we found that type  
413 1-3 OM particles not only occurred in 70% aerosol particles but also were  
414 concentrated on fine particles (< 1 μm) (Fig. 7a). Therefore, these OM-containing  
415 particles in the frequent L&M hazes could pose a threat to human health for one long  
416 period throughout the winter. Although haze episodes commonly occur in the NCP,  
417 the morphology and mixing state of OM particles in the wintertime L&M hazes are  
418 largely different from those in the more severe wintertime hazes and from those in  
419 L&M hazes in other seasons (Li and Shao, 2009; Li and Shao, 2010). Therefore, these  
420 microscopic observations provide important information on the different haze  
421 formations and evaluate their atmospheric impacts on climate and human health.



422

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646

647



## Figure Captions

648

649 **Figure 1.** Cluster-means of backward trajectories at S1 (urban Jinan), S2 (Mt. Tai top),  
650 and S3 (Changdao island) sites during 13-23 December, 2014. Air masses on clear  
651 days normally are from the northwest and air masses on haze days are from the west  
652 or southwest. MODIS images show the layer during the light and moderate regional  
653 hazes over the NCP.

654 **Figure 2.** NanoSIMS-based ion intensity mappings of  $^{12}\text{C}^{14}\text{N}^-$ ,  $^{14}\text{N}^{16}\text{O}_2^-$ ,  $^{16}\text{O}^-$ , and  $^{32}\text{S}^-$   
655 from a near-spherical OM particle.

656 **Figure 3.** Typical TEM images of different types of OM particles. (a) Type 1:  
657 spherical shape; (b) type 2: near-spherical shape; (c) type 3: irregular shape; (d) type 4:  
658 domelike OM (droplet-like particle); (e) type 5: dispersed-OM; (f) type 6:  
659 OM-coating

660 **Figure 4.** Number fraction of OM types in particulate matter during light and  
661 moderate haze episodes from 13 to 23 December 2014. The 2562 OM particles were  
662 analyzed by TEM-EDX.

663 **Figure 5.** Typical TEM images of OM internally mixed particles (a) a spherical OM  
664 particle attached to a soot particle; (b) a near-spherical OM particle attached to a  
665 mineral particle; (c) fly ash particles attached to a near-spherical OM particle; (d) OM  
666 mixed with sulfate (K)-fly ash particle and its sublimed particle under strong electron  
667 beam; (e) OM as a coating mixed with a sulfate (K, Na) particle. The element  
668 compositions in OM, mineral, fly ash, and sulfate particles were measured by the  
669 TEM/EDX.

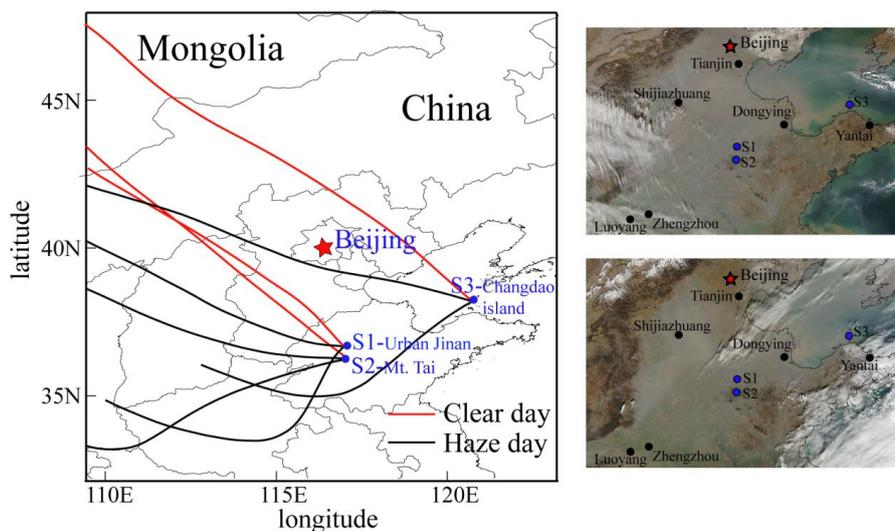
670 **Figure 6.** Number fraction of OM internally mixed particles at (a) S1 site, (b) S2 site,  
671 and (c) S3 site. The number of analyzed particles in different size ranges is shown  
672 above each column.

673 **Figure 7.** Size distributions of OM-containing particles and OM particles during the  
674 same hazes. (a) Size distributions of 1088, 1341 and 1136 OM-containing particles  
675 collected at S1, S2, and S3. (b) Size distributions of 269, 173, and 297 type 1-3 OM  
676 particles collected at S1, S2, and S3.

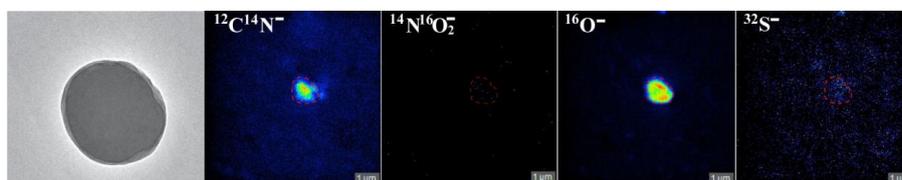


677 **Figure 8.** Triangular diagram of weight ratios of Si-O-C based on TEM/EDX data. 34  
678 OM particles from coal combustion and 28 OM particles from corn stalks combustion,  
679 and 281 OM particles produced from the haze samples in this study. The three lines  
680 represent the links of the ratios of Si/O and Si/C in haze, corn stalks and coal  
681 combustion samples, respectively.

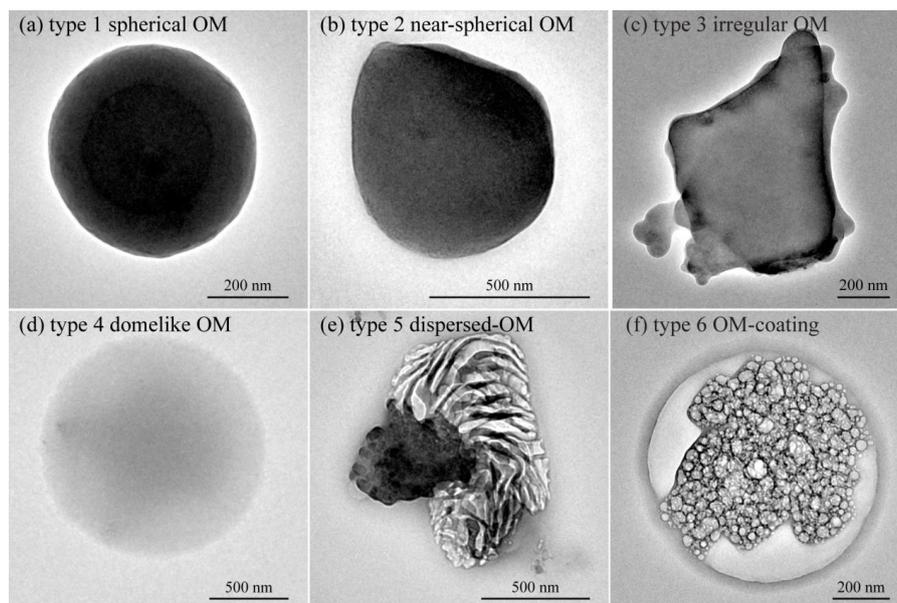
682 **Figure 9.** The relationship between the size of individual particles and their sulfate  
683 cores based on 366 OM-coating particles at S1, S3, and S3 sites. The smaller slope  
684 represents the thicker OM coating. The number fractions of OM coating to  
685 OM-containing particles at three sampling sites are shown in the pie charts.



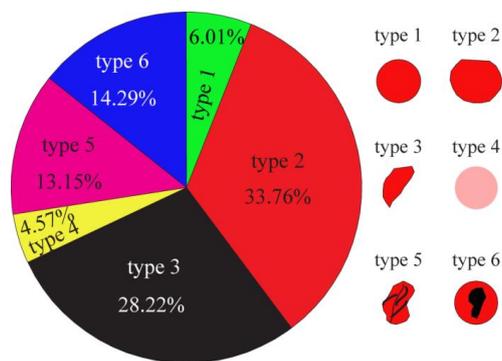
**Figure 1.** Cluster-means of backward trajectories at S1 (urban Jinan), S2 (Mt. Tai top), and S3 (Changdao island) sites during 13-23 December, 2014. Air masses on clear days normally are from the northwest and air masses on haze days are from the west or southwest. MODIS images show the layer during the light and moderate regional hazes over the NCP.



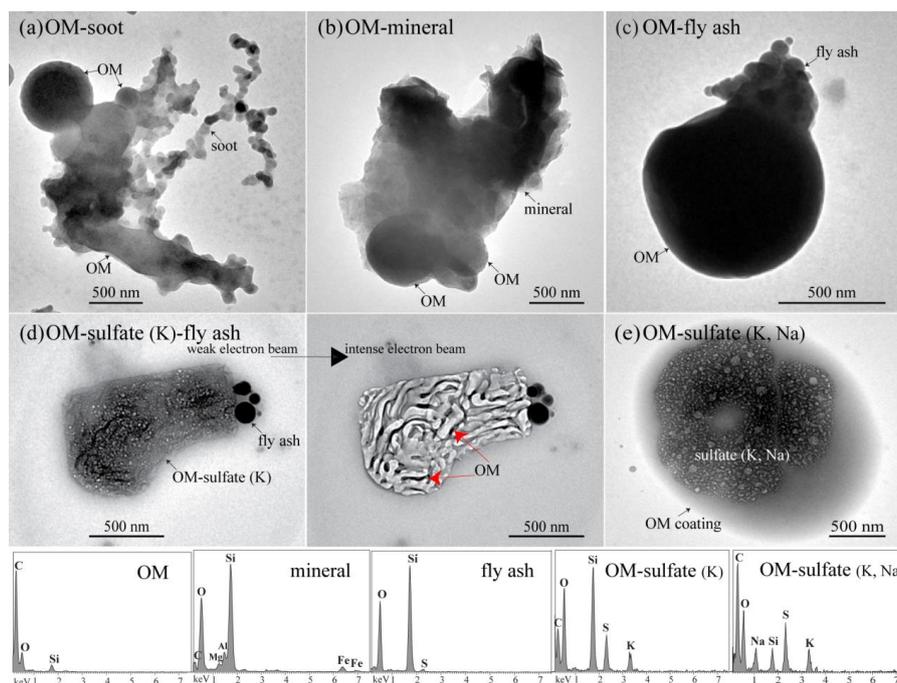
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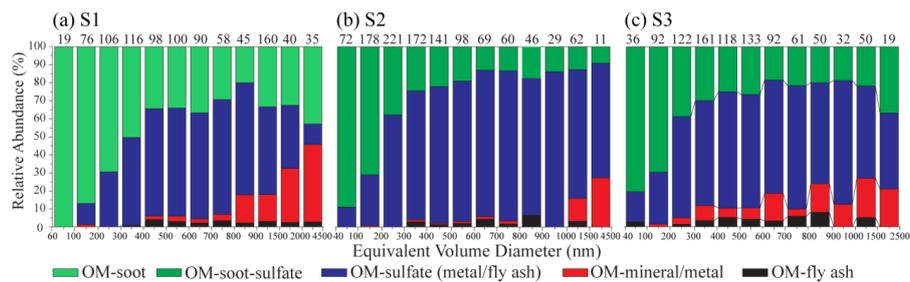
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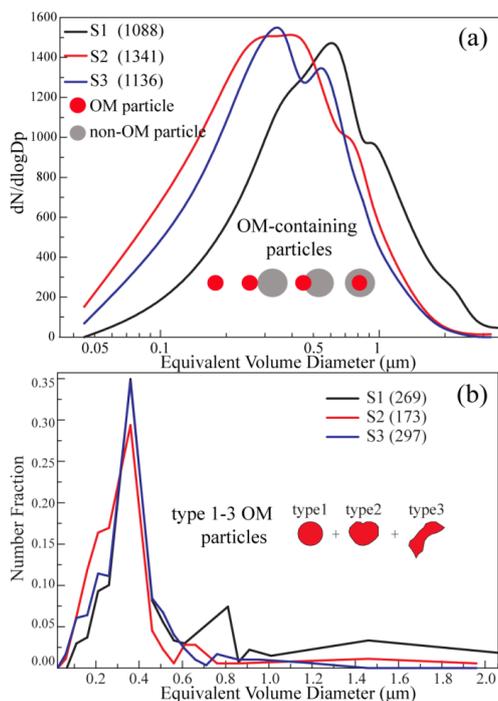


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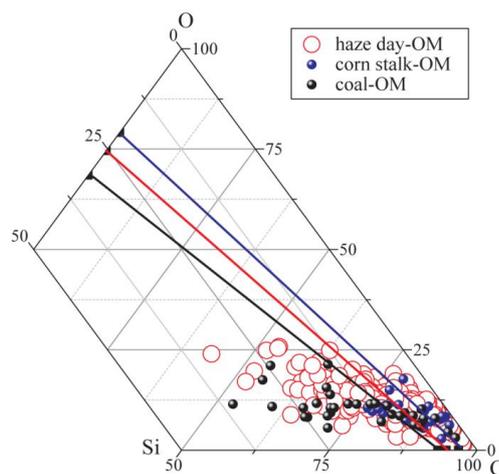




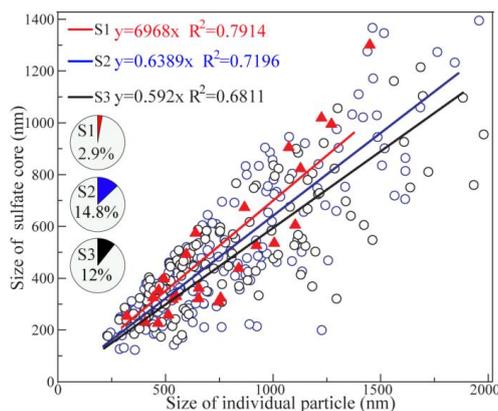
**Figure 6.** Number fraction of OM internally mixed particles at (a) S1 site, (b) S2 site, and (c) S3 site. The number of analyzed particles in different size ranges is shown above each column.



**Figure 7.** Size distributions of OM-containing particles and OM particles during the same hazes. (a) Size distributions of 1088, 1341 and 1136 OM-containing particles collected at S1, S2, and S3. (b) Size distributions of 269, 173, and 297 type 1-3 OM particles collected at S1, S2, and S3.



**Figure 8.** Triangular diagram of weight ratios of Si-O-C based on TEM/EDX data. 34 OM particles from coal combustion and 28 OM particles from corn stalks combustion, and 281 OM particles produced from the haze samples in this study. The three lines represent the links of the ratios of Si/O and Si/C in haze, corn stalks and coal combustion samples, respectively.



**Figure 9.** The relationship between the size of individual particles and their sulfate cores based on 366 OM-coating particles at S1, S3, and S3 sites. The smaller slope represents the thicker OM coating. The number fractions of OM coating to OM-containing particles at three sampling sites are shown in the pie charts.



**Table 1.** Average size, number, sphericity, and aspect ratio for different OM types at the three sampling sites.

Sampling site	Type	Average Size (nm)	Number	Average Sph (min, max); ( $\pm$ stdev.)	Average AR (min, max); ( $\pm$ stdev.)
S1	Type1 spherical	407.8	18	0.9664 (0.8, 1.0); ( $\pm$ 0.05)	1.0223 (1.0, 1.16); ( $\pm$ 0.03)
	Type2 near spherical	348.5	79	0.8172 (0.50, 1); ( $\pm$ 0.11)	1.1275 (1.0, 1.41); ( $\pm$ 0.08)
	Type3 irregular	538.68	151	0.5046 (0.07,0.95); ( $\pm$ 0.21)	1.4764 (1.04, 3.57); ( $\pm$ 0.41)
	Type4 domelike	758.76	23	0.6374 (0.29, 0.9); ( $\pm$ 0.21)	1.3958 (1.08, 1.89); ( $\pm$ 0.26)
	Type5 dispersed-OM	750.8	64	0.4940 (0.1, 0.92); ( $\pm$ 0.21)	1.4646 (1.06, 2.85); ( $\pm$ 0.40)
	Type6 OM-coating	672.18	12	0.9251(0.45, 1); ( $\pm$ 0.13)	1.0627 (1, 1.45); ( $\pm$ 0.11)
S2	Type1 spherical	282.15	22	0.9628 (0.76, 1.0); ( $\pm$ 0.06)	1.0281 (1.00, 1.25); ( $\pm$ 0.05)
	Type2 near spherical	323.15	68	0.8254 (0.53, 1.0); ( $\pm$ 0.12)	1.1197 (1.00, 1.37); ( $\pm$ 0.09)
	Type3 irregular	399.26	62	0.5746 ((0.06,0.96); ( $\pm$ 0.24)	1.4102 (1.04, 2.79); ( $\pm$ 0.42)
	Type4-domelike	511.35	25	0.7341 (0.15, 1.0); ( $\pm$ 0.29)	1.2433 (1.00, 2.28); ( $\pm$ 0.36)
	Type5-dispersed-OM	846.99	34	0.5773 (0.1, 0.9); ( $\pm$ 0.25)	1.4218 (1.08, 2.71) ( $\pm$ 0.45)
	Type6-OM-coating	775.53	66	0.8791 (0.26, 1) ( $\pm$ 0.22)	1.0797 (1, 1.57) ( $\pm$ 0.12)
S3	Type1 spherical	391.5	27	0.9901 (0.92, 1.0); ( $\pm$ 0.02)	1.0065 (1.0, 1.06); ( $\pm$ 0.02)
	Type2 near spherical	373.09	122	0.8341 (0.33, 1.0); ( $\pm$ 0.13)	1.12 (1,1.51); ( $\pm$ 0.10)
	Type3 irregular	374.85	117	0.5675 (0.09, 0.96); ( $\pm$ 0.21)	1.392 (1.04, 2.82); ( $\pm$ 0.34)
	Type4 domelike	705.94	14	0.6325 (0.22, 1); ( $\pm$ 0.23)	1.3519 (1, 1.857); ( $\pm$ 0.22)
	Type5 dispersed-OM	574.62	35	0.5495 (0.4,0.65); ( $\pm$ 0.19)	1.3475 (1.21, 1.51); ( $\pm$ 0.31)
	Type6 OM-coating	828.35	28	0.9291 (0.36,1); ( $\pm$ 0.30)	1.0591 (1.0, 1.60); ( $\pm$ 0.14)