

# 1 **Measurements of biogenic volatile organic compounds at a** 2 **grazed savannah-grassland-agriculture landscape in South** 3 **Africa**

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## 17 18 **Abstract**

19 Biogenic volatile organic compounds (BVOCs) are important role players in the chemistry of  
20 the troposphere, especially in the formation of tropospheric ozone (O<sub>3</sub>) and secondary organic  
21 aerosols (SOA). Ecosystems produce and emit a large number of BVOCs. It is estimated on a  
22 global scale that approximately 90 % of annual BVOC emissions are from terrestrial sources.  
23 In this study, measurements of BVOCs were conducted at the Welgegund measurement station  
24 (South Africa), which is considered to be a regionally representative background site situated  
25 in savannah grassland. Very few BVOC measurements exist for grassland savannah and results  
26 presented in this study are the most extensive for this type of landscape. Samples were collected  
27 twice a week for two hours during daytime and two hours during night-time through two long-  
28 term sampling campaigns from February 2011 to February 2012 and from December 2013 to

1 February 2015. Individual BVOCs were identified and quantified using a thermal desorption  
2 instrument, connected to a gas chromatograph and a mass selective detector. The annual  
3 median concentrations of isoprene, 2-methyl-3-butene-2-ol (MBO), monoterpenes and  
4 sesquiterpenes (SQT) during the first campaign were 14, 7, 120 and 8 pptv, respectively, and  
5 14, 4, 83 and 4 pptv, respectively, during the second campaign. The sum of the concentrations  
6 of the monoterpenes were at least an order of magnitude higher than the concentrations of other  
7 BVOC species during both sampling campaigns, with  $\alpha$ -pinene being the most abundant  
8 species. The highest BVOC concentrations were observed during the wet season and elevated  
9 soil moisture was associated with increased BVOC concentrations. However, comparisons  
10 with measurements conducted at other landscapes in southern Africa and the rest of the world  
11 that have more woody vegetation indicated that BVOC concentrations were, in general,  
12 significantly lower for the grassland savannah. Furthermore, BVOC concentrations were an  
13 order of magnitude lower compared to total aromatic concentrations measured at Welgegund.  
14 An analysis of concentrations by wind direction indicated that isoprene concentrations were  
15 higher from the western sector that is considered to be a relatively clean regional background  
16 region with no large anthropogenic point sources, while wind direction did not indicate any  
17 significant differences in the concentrations of the other BVOC species. Statistical analysis  
18 indicated that soil moisture had the most significant impact on atmospheric levels of MBO,  
19 monoterpenes and SQT concentrations, while temperature had the greatest influence on  
20 isoprene levels. The combined O<sub>3</sub> formation potentials of all the BVOCs measured calculated  
21 with MIR coefficients during the first and second campaign were 1162 and 1022 pptv,  
22 respectively.  $\alpha$ -Pinene and limonene had the highest reaction rates with O<sub>3</sub>, while isoprene  
23 exhibited relatively small contributions to O<sub>3</sub> depletion. Limonene,  $\alpha$ -pinene and terpinolene  
24 had the largest contributions to the OH-reactivity of BVOCs measured at Welgegund for all of  
25 the months during both sampling campaigns.

## 26 **1 Introduction**

27 Ecosystems produce and emit a large number of biogenic volatile organic compounds (BVOCs)  
28 that are involved in plant growth and reproduction. These species also act as defensive  
29 compounds, e.g. enhancing tolerance to heat and oxidative stress (Sharkey and Yeh, 2001;  
30 Loreto and Schnitzler, 2010), preventing the colonisation of pathogens after wounding, and  
31 deterring insects or recruiting natural enemies of herbivores (Holopainen and Gershenzon,  
32 2010). The BVOC production rate in an ecosystem depends on several physical (e.g.

1 temperature, precipitation, moisture, solar radiation and CO<sub>2</sub> concentration) and biological  
2 parameters (e.g. plant species and the associated emission capacity, phenology, biotic and  
3 abiotic stresses, attraction of pollinators) (Blande et al., 2014; Fuentes et al., 2000; Kesselmeier  
4 and Staudt, 1999; Sharkey and Yeh, 2001), with typically 0.2 to 10 % of the carbon uptake  
5 during photosynthesis being converted to BVOCs (Kesselmeier et al., 2002). It is estimated  
6 that, on a global scale, approximately 90 % of annual BVOC emissions are from vegetation/  
7 terrestrial sources (~1000 Tg year<sup>-1</sup>) (Guenther et al., 2012).

8 BVOCs can contribute significantly to the carbon balance in certain ecosystems (Kesselmeier  
9 et al., 2002; Malhi, 2002). BVOC concentrations in ambient air depend on several factors,  
10 which include emission rates from vegetation, atmospheric transport and mixing, as well as the  
11 chemical composition and oxidative state of the atmosphere, which determines the sink of these  
12 species. BVOCs are important in the formation of tropospheric ozone (O<sub>3</sub>) and secondary  
13 organic aerosols (SOA). BVOCs in the troposphere react with the major oxidants in the  
14 atmosphere, which include tropospheric O<sub>3</sub>, hydroxyl radicals (<sup>•</sup>OH, referred to, from here on,  
15 as OH for simplicity) and nitrate radicals (NO<sub>3</sub><sup>•</sup>, referred to, from here on, as NO<sub>3</sub> for simplicity)  
16 (Atkinson and Arey, 2003a). These oxidants strongly affect the concentrations of atmospheric  
17 BVOCs (Lelieveld et al., 2008; Di Carlo et al., 2004). BVOCs are also crucial in the formation  
18 of the stabilised Criegee intermediate – a carbonyl oxide with two free-radical sites – or its  
19 derivative (Mauldin III et al., 2012; Welz et al., 2012), which also contributes to atmospheric  
20 oxidation. A complex range of reaction products are formed from atmospheric BVOCs,  
21 including less volatile oxygenated compounds that condense to form aerosol particles.

22 Various studies have indicated the link between BVOCs and the formation of SOA (Vakkari et  
23 al., 2015; Andreae and Crutzen, 1997; Ehn et al., 2014), while the influence of BVOCs on the  
24 growth of newly formed aerosol particles has also been indicated (Kulmala et al., 2004; Tunved  
25 et al., 2006). However, there are many uncertainties associated with the exact chemical  
26 reactions and physical processes involved in SOA formation and aerosol particle growth, which  
27 largely depends on regional emissions and atmospheric processes (Kulmala et al., 2013; Ehn et  
28 al., 2014). Vakkari et al. (2015) indicated the importance of VOCs for new particle formation  
29 and growth in clean background air in South Africa. Therefore, it is essential to understand the  
30 sources, transport and transformations of these compounds for air quality management and  
31 climate change-related studies, as well as for the modelling of atmospheric chemistry at global,

1 regional and local scales (Laothawornkitkul et al., 2009; Peñuelas and Staudt, 2010; Peñuelas  
2 and Llusà, 2003).

3 Long-term ambient BVOC measurements to establish seasonal cycles have been conducted  
4 extensively in several regions, which include boreal forest (Hakola et al., 2009; Hakola et al.,  
5 2000; Rinne et al., 2000; Rinne et al., 2005; Rantala et al., 2015; Räisänen et al., 2009;  
6 Eerdekens et al., 2009; Lappalainen et al., 2009), hemiboreal mixed forest (Noe et al., 2012),  
7 temperate (Spirig et al., 2005; Stroud et al., 2005; Fuentes et al., 2007; Mielke et al., 2010),  
8 Mediterranean (Davison et al., 2009; Harrison et al., 2001) and tropical (Rinne et al., 2002)  
9 ecosystems. Shorter campaigns have also been conducted in Western and Central Africa, which  
10 include several different studies in the framework of the African Monsoon Multidisciplinary  
11 Analyses (AMMA) (Grant et al., 2008; Saxton et al., 2007) and the EXPeriment for the  
12 REgional Sources and Sinks of Oxidants (EXPRESSO) (Serca et al., 2001). Zunckel et al.  
13 (2007) and references therein indicated that limited research has been conducted on BVOC  
14 emissions in southern Africa, which consisted mainly of short campaigns measuring BVOC  
15 emission rates. Considering that BVOC emissions on a global scale are considered to be  
16 significantly higher (ca. 10 times) than the emission of anthropogenic VOCs, it is very  
17 important that longer-term BVOC measurements are conducted in southern Africa.  
18 Furthermore, a large part of the land cover in South Africa consists of a grassland bioregion, as  
19 indicated in Figure 1. Although it is considered that grasslands cover approximately one quarter  
20 of the Earth's land surface, relatively few studies have been conducted on BVOC emissions  
21 from grasslands, while there are no long-term BVOC studies reported for these landscapes  
22 (Bamberger et al., 2011; Ruuskanen et al., 2011; Wang et al., 2012). Therefore, the aim of this  
23 study was to quantify the ambient BVOC concentrations over different seasons at a regional  
24 background site in South Africa. In addition, the objective was also to characterise their  
25 seasonal patterns, as well as to relate BVOC concentrations measured in southern Africa to  
26 levels in other regions in the world. According to the knowledge of the authors, this is the first  
27 record of ambient BVOC concentrations covering a full seasonal cycle in southern Africa and  
28 for a grassland bioregion anywhere in the world.

29

30 **Insert Figure 1**

31

## 1 **2 Measurement location and methods**

### 2 **2.1 Site description**

3 Measurements were conducted at the Welgegund measurement station (26.57°S, 26.94°E, 1480  
4 m a.s.l.) (Welgegund measurement station, 2016), which is located on the property of a  
5 commercial maize and cattle farmer approximately 100 km west of Johannesburg, as indicated  
6 in Figure 1. Welgegund is a regional background station with no pollutant sources in close  
7 proximity. The distances to the nearest blacktop road and nearest town are approximately 10  
8 and 30 km, respectively. Welgegund is, however, affected by the major anthropogenic source  
9 regions in the north-eastern interior of South Africa (as indicated by the major large point  
10 sources in Figure 1), which also include the Johannesburg-Pretoria conurbation (Tiitta, et al.,  
11 2014). From Figure 1, it is also evident that the western sector contains no major anthropogenic  
12 point sources and can therefore be considered to be representative of a relatively clean regional  
13 background.

14 Welgegund is geographically located within the South African Highveld, which is characterised  
15 by two distinct seasonal periods, i.e. a dry season from May to September that predominantly  
16 coincides with winter (June to August), and a wet season during the warmer months from  
17 October to April. The dry period is characterised by low relative humidity, while the wet season  
18 is associated with higher relative humidity and frequent rains that predominantly occur in the  
19 form of thunderstorms. The mean annual precipitation is approximately 500 mm with  
20 approximately >80 % of rain events occurring during the wet season. During the sampling  
21 period, the coldest temperature recorded in winter at Welgegund was -5.1 °C in June 2011,  
22 while the highest temperature recorded in summer was +35.6 °C in October 2011. The mean  
23 maximum temperature ranges between 16 and 32 °C, while the mean minimum temperature  
24 ranges between 0 and 15 °C. Winters are also characterised by frequent and severe frost days  
25 (26-37 days per year) (Mucina and Rutherford, 2006).

### 26 **2.2 Vegetation**

27 The Welgegund measurement station is located in the Grassland Biome (Figure 1), which  
28 covers 28 % of South Africa's land surface (Mucina and Rutherford, 2006). This biome has  
29 been significantly transformed, primarily as a result of cultivation, plantation forestry,  
30 urbanisation and mining (Daemane et al., 2010 and references therein). It has also been severely

1 degraded by erosion and agricultural development. The station is situated within Vaal-Vet  
2 Sandy Grassland with Andesite Mountain Bushveld of the Savannah Biome prominent on  
3 nearby ridges. At present, only 0.3 % of the Vaal-Vet Sandy Grassland is statutorily conserved,  
4 while the rest is mostly used for grazing and crop production. In Figure 2, a land cover map  
5 within a 60 km radius from Welgegund is presented indicating the extent of cultivation in this  
6 region. The land cover survey was performed within a region that was estimated to represent  
7 the BVOC footprint at Welgegund, which was calculated from typical atmospheric lifetimes  
8 (Table 1) of the species measured and the general wind speed(s) (Figure 3) at Welgegund. The  
9 immediate area surrounding Welgegund is grazed by livestock, with the remaining area covered  
10 by crop fields (mostly maize and to a lesser degree sunflower). In the demarcated 60 km radius,  
11 a further three vegetation units of the Dry Highveld Grassland Bioregion (Grassland Biome)  
12 and another two of the Central Bushveld Bioregion (Savannah Biome) are also present. In  
13 addition, alluvial vegetation is found associated with major rivers and inland saline vegetation  
14 in scattered salt pans.

15

## 16 **Insert Figure 2**

17

18 The study area comprises a highly variable landscape with scattered hills and sloping, slightly  
19 irregular, undulating plains, which are dissected by prominent rocky ridges. Soil in the  
20 catchment area is heterogeneous and rocky, varying from sandy to clayey depending on the  
21 underlying rock types, such as andesite, chert, dolomite, mudstone, quartzite, sandstone and  
22 shale.

23 Land use within the surrounding area is divided into six major land cover types, i.e. cultivated  
24 land, grasslands, mountainous areas, plantations, urban areas and water bodies, as indicated in  
25 Figure 2. Mountainous areas, grassland and water bodies (riparian areas) comprised many  
26 different vegetation units. The other homogenous areas were anthropogenically altered and no  
27 longer representative of the surrounding natural vegetation. The study area is characterised by  
28 a grassland-woodland vegetation complex, dominated by various grass and woody species, and  
29 recognised by the presence of non-native species in altered environments.

30 The most dominant woody species of the entire study area include the trees *Celtis africana*,  
31 *Searsia pyroides*, *Vachellia karroo* and *Ziziphus mucronata*, and the thorny shrub *Asparagus*

1 *laricinus*. Tree diversity increases where there are patches of deep sand, characterised by  
2 *Gymnosporia buxifolia* and *Vachellia erioloba*, or in mountainous areas, where *Euclea*  
3 *undulata*, *Grewia flava* and *Senegallia caffra* become most prominent. Woody vegetation  
4 occurs sparsely in grasslands and when present is found on isolated ridges, including the small  
5 trees *Pavetta zeyheri*, *Vangueria infausta* and *Zanthoxylum capense*. In anthropogenically  
6 altered environments, native species decrease and introduced species dominate, such as  
7 *Eucalyptus camaldulensis*, *Pinus roxburghiana* and *Populus canescens* in plantations; *Gleditsia*  
8 *triacanthos*, *Pyracantha coccinea* and *Salix babylonica* along rivers and water bodies; and  
9 *Celtis sinensis*, *Melia azedarach* and *Robinia pseudoacacia* in the urban footprint.

10 The most dominant species of the grass sward in the entire study area include *Cynodon*  
11 *dactylon*, *Eragrostis chloromelas*, *Heteropogon contortus*, *Setaria sphacelata* and *Themeda*  
12 *triandra*. The dry, western grassland (Vaal-Vet Sandy Grassland specifically) is characterised  
13 by *Antheophora pubescens*, *Cymbopogon caesius*, *Digitaria argyrograpta*, *Elionurus muticus*  
14 and *Eragrostis lehmanniana*, and the moist Rand Highveld Grassland in the south-east by  
15 *Ctenium concinnum*, *Digitaria monodactyla*, *Monocymbium ceresiforme*, *Panicum natalense*  
16 and *Trachypogon spicatus*. The north-eastern parts of the study area on dolomite are dominated  
17 by *Brachiaria serrata*, *Digitaria tricholaenoides*, *Eragrostis racemosa* and *Loudetia simplex*.

## 18 **2.3 Measurement methods**

### 19 **2.3.1 BVOC measurements and analysis**

20 BVOC measurements were conducted for a period of more than two years through a 13-month  
21 sampling campaign from February 2011 to February 2012 and a 15-month sampling campaign  
22 from December 2013 to February 2015. Samples were collected twice a week for two hours  
23 during daytime (11:00 to 13:00 local time, LT) and two hours during night-time (23:00 to 1:00  
24 LT) on Tuesdays and Saturdays. Several previous studies have demonstrated that the maximum  
25 emissions of isoprene and monoterpenes from vegetation occur around midday (Fuentes et al.,  
26 2000; Kuhn et al., 2002). Understandably, the chosen sampling schedule, i.e. same days each  
27 week and same hours of the day, was prone to some bias. As mentioned by Jaars et al. (2014),  
28 considering the distance of the sampling site from the nearest town and logistical limitations  
29 during the sampling campaigns, the sampling schedule applied was the most feasible option  
30 that enabled the collection of data for more than two years. VOCs were sampled at a height of  
31 2 m above ground level, with a 1.75 m long inlet. The first 1.25 m of the inlet was a stainless

1 steel tube (grade 304 or 316) and the second 0.5 m was Teflon. To prevent the degradation of  
2 BVOC by O<sub>3</sub>, the stainless steel part of the inlet was heated to 120 °C using heating cables and  
3 thermostats (Thermonic), thereby removing ozone from the sample stream (Hellén et al.,  
4 2012a). At regular intervals, the efficiency of this O<sub>3</sub> removal was verified with an O<sub>3</sub> monitor,  
5 which indicated that O<sub>3</sub> concentrations decreased from median values  $\geq 30$  ppb to  $< 2$  ppb (Jaars  
6 et al., 2014).

7 VOCs were collected with stainless steel adsorbent tubes (6.3 mm ED x 90 mm, 5.5 mm ID)  
8 packed with Tenax-TA and Carbotrap-B by using a constant flow type automated  
9 programmable sampler. A detailed description of the sampling procedure is presented by Jaars  
10 et al. (2014). In short, the flow rate of the pump was set at between 100 and 110 ml min<sup>-1</sup>  
11 throughout the campaigns and was calibrated each week. Prior to sampling, all adsorbent tubes  
12 were tested for leaks and preconditioned with helium for 30 minutes at 350 °C at a flow of 40  
13 ml min<sup>-1</sup>. The adsorbent tubes were removed from the sampler once a week and closed off with  
14 Swagelok® caps. Each tube was stored in a container for transport to the laboratory, where the  
15 adsorbent tubes were stored in a freezer for two to four weeks prior to analysis.

16 Individual BVOCs were identified and quantified using a thermal desorption instrument  
17 (Perkin-Elmer TurboMatrix™ 650, Waltham, USA) connected to a gas chromatograph  
18 (Perkin-Elmer® Clarus® 600, Waltham, USA) with a DB-5MS (60 m, 0.25 mm, 1 µm) column  
19 and a mass selective detector (Perkin-Elmer® Clarus® 600T, Waltham, USA). Samples were  
20 analysed using the selected ion mode (SIM). A five-point calibration was performed by using  
21 liquid standards in methanol solutions. Standard solutions were injected onto adsorbent tubes  
22 that were flushed with helium at a flow of 100 ml min<sup>-1</sup> for 10 min in order to remove methanol.  
23 BVOCs quantified for the two campaigns included isoprene with method detection limit (MDL)  
24 between 1.2 and 2.4 pptv and 2-methyl-3-butene-2-ol (MBO) with MDL  
25 between 0.9 and 1.4 pptv. The monoterpenes ( $\alpha$ -pinene, camphene,  $\beta$ -pinene,  $\Delta^3$ -carene,  
26 limonene, 1,8-cineol, terpinolene, nopinone and bornylacetate) and monoterpene-related  
27 BVOCs (p-cymene, 4-allylanisole and 4-acetyl-1-methylcyclohexene (AMCH)) MDL was  
28 between 0.6 and 1.6 pptv. The term “monoterpene(s)” used in the discussions in subsequent  
29 sections in the manuscript refers to both the monoterpene and monoterpene-related BVOCs.  
30 The sesquiterpenes (SQT) (longicyclene, iso-longifolene, aromadendrene,  $\alpha$ -humulene and  
31 alloaromadendrene) MDL was ~0.6 pptv. Since the analytical system did not separate myrcene  
32 and  $\beta$ -pinene,  $\beta$ -pinene concentrations determined were the sum of these two species. VOC



1 concentrations were field and lab blank corrected. When monthly median BVOC  
2 concentrations were calculated, sample concentrations below the method detection limit (MDL)  
3 were replaced with  $\frac{1}{2}$ MDL.

### 4 **2.3.2 Ancillary measurements**

5 Ancillary measurements continuously performed at the Welgegund station were used to  
6 interpret the measured BVOC concentrations. General meteorological parameters, i.e.  
7 temperature (T), relative humidity (RH), wind speed and direction, and precipitation were  
8 measured. Soil temperature and moisture at different depths (5 and 20 cm) were measured with  
9 a PT-100 and Theta probe ML2x (Delta-T), respectively. Additional soil moisture information  
10 was obtained with a 100 cm PR2 soil moisture profile probe (Delta-T). Direct photosynthetic  
11 photon flux density (PPFD) between 400 and 700 nm was measured with a Kipp & Zonen  
12 pyranometer (CMP 3 pyranometer, ISO 9060:1990 Second Class).

13 Trace gas measurements were performed utilising a Thermo-Electron 43S sulphur dioxide  
14 ( $\text{SO}_2$ ) analyser (Thermo Fisher Scientific Inc., Yokohama-shi, Japan), a Teledyne 200AU  
15 nitrogen oxide ( $\text{NO}_x$ ) analyser (Advanced Pollution Instrumentation Inc., San Diego, Cam  
16 USA), an Environment SA 41M  $\text{O}_3$  analyser (Environment SA, Poissy, France) and a Horiba  
17 APMA-360 carbon monoxide (CO) analyser (Horiba, Kyoto, Japan). The net ecosystem  
18 exchange (NEE) of carbon dioxide ( $\text{CO}_2$ ) was measured with the eddy covariance method with  
19 a Licor 7000 closed path infrared gas analyser (IRGA) and a three-dimensional Metek sonic  
20 anemometer at a height of 9 m, which is well above the average tree height of 2.5 m (Räsänen  
21 et al., 2016). A more detailed description of additional parameters monitored at Welgegund is  
22 given by Beukes et al. (2015).

### 23 **2.3.3 Lifetime of BVOCs**

24 In Table 1, the atmospheric lifetimes ( $\tau$ ) of BVOCs measured in this study calculated from OH-  
25 and  $\text{O}_3$  reactivity are reported. BVOC lifetimes according to  $\text{O}_3$  reactivity were calculated with  
26 Eq. (1):

$$27 \quad \tau = \tau_{\text{O}_3} = \frac{1}{k_{\text{O}_3, [\text{O}_3]}} \quad (1)$$

28 where  $[\text{O}_3]$  is the annual average  $\text{O}_3$  concentration (ca. 36 ppbv) measured during the two  
29 campaigns at Welgegund and  $k_{\text{O}_3}$  the reaction rate constant for the reaction between a specific

1 BVOC and O<sub>3</sub>. Since direct OH reactivity measurements were not available, the average  
2 concentration of OH radicals ([OH]) (ca. 0.04 pptv) reported by Ciccioli et al. (2014) was used  
3 in the calculations, using Eq. (2):

$$4 \quad \tau = \tau_{\text{OH}} = \frac{1}{k_{\text{OH},[\text{OH}]}} \quad (2)$$

5 where  $k_{\text{OH}}$  is the reaction rate constant for the reaction between a specific BVOC and OH.

6

7 **Insert Table 1**

8

### 9 **3 Results and discussion**

#### 10 **3.1 Meteorological conditions during the measurement campaigns**

11 Local meteorological influences on the measured BVOC concentrations are likely to be more  
12 significant than regional impacts of air masses due to the short lifetimes associated with  
13 atmospheric BVOCs (Table 1). Therefore, BVOC concentrations were only interpreted in  
14 terms of local meteorological patterns and no back trajectory analyses were employed. In  
15 Figure 3, the monthly medians of the meteorological parameters – precipitation, T, RH, wind  
16 speed and -direction, and soil moisture depth (5 and 20 cm) – measured at Welgedund during  
17 each of the two sampling campaigns are presented. From Figure 3a and b, the wet season  
18 (October to April) associated with warmer months and the dry season (May to September)  
19 associated with colder months as discussed in section 2.1 are evident. Rainfall in this region of  
20 South Africa is typically characterised by relatively large inter-annual variability (Conradie et  
21 al., 2016). The monthly median temperatures for the periods during which samples were  
22 collected ranged between 8.8 and 13 °C in winter and 19.7 and 24.9 °C in summer (Figure 3b).  
23 During the warmer months, temperatures up to 30 °C and higher were reached frequently.  
24 During the wet season, the monthly median RH ranged between 30 (with the onset of the wet  
25 season) and 80 % (at the end of the wet season), while the RH ranged between 20 and 50 %  
26 during the dry season (Figure 3c). The highest monthly median wind speeds occurred during  
27 the warmer months (Figure 3d) when unstable meteorological conditions are prevalent in the  
28 interior of South Africa (Tyson et al., 1996). The seasonal variations of wind direction during  
29 the two sampling campaigns (Figure 3e) indicated that the prevailing wind direction was from

1 the northern to eastern sector, which agrees with the back trajectory analysis performed for the  
2 first sampling period at Welgegund by Jaars et al. (2014). Soil moisture measurements  
3 mimicked the seasonal precipitation pattern, i.e. higher soil moisture associated with the wet  
4 season (Figure 3f and 3g). The soil moisture measurements conducted from January to August  
5 at a depth of 20 cm were significantly higher during the first sampling campaign. During  
6 December 2010 and January 2011, prior to the first sampling campaign, precipitation (Figure  
7 3a) was clearly higher than during the second campaign, i.e. December 2013 to January 2014.  
8 Subsequently, the soil moisture measured at 20 cm (Figure 3g) was clearly higher during the  
9 first sampling campaign than during the second campaign from the beginning of the campaign  
10 until the middle of the dry season.

11

### 12 **Insert Figure 3**

13

14 Figure 4 presents micrometeorological CO<sub>2</sub> flux measurements at Welgegund, which indicate  
15 typical changes in the seasonal uptake of CO<sub>2</sub> by vegetation. In addition, the Photosynthetic  
16 Photon Flux Density (PPFD) is also indicated with a colour bar. Negative values (downward  
17 CO<sub>2</sub> flux) indicate the net uptake of CO<sub>2</sub> by vegetation, with the gross primary production  
18 (GPP) exceeding the total respiration. Positive values indicate the emission of CO<sub>2</sub> by the  
19 vegetation. A period of an approximately 0 (small positive) net CO<sub>2</sub> flux is observed in the  
20 winter months that extend until September, which can be attributed to decreased microbial  
21 activity associated with lower temperatures, low rainfall and most of the vegetation losing their  
22 leaves. The net ecosystem exchange (NEE) at full light (maximum downward flux) increases  
23 gradually until February in response to the increases of the photochemical efficiency of CO<sub>2</sub>  
24 assimilation in the vegetation surrounding the site and the solar elevation angle. The daily  
25 maximum NEE starts to decrease in March/April when the solar elevation angle declines and  
26 soil moisture drops.

27

### 28 **Insert Figure 4**

29

## 1 **3.2 Contextualising BVOC concentrations measured at Welgegund**

2 In Table 2, the median (mean) and inter-quartile range (IQR, 25<sup>th</sup> to 75<sup>th</sup>) concentrations, as  
3 well as the median (mean) daytime to night-time concentration ratios of the BVOC species  
4 determined during the two sampling campaigns at Welgegund are presented. It is evident from  
5 the median (mean) daytime to night-time concentration ratios that there were not significant  
6 differences in levels of most of the BVOCs measured during daytime and night-time at  
7 Welgegund, with the exception of isoprene measured during the first sampling campaign, as  
8 well as the monoterpenes terpinolene and bornylacetate, and the SQT aromadendrene measured  
9 during the second sampling campaign. Isoprene levels during the first sampling campaign were  
10 approximately two times higher during daytime, which reflect the light dependency usually  
11 associated with isoprene emissions. However, daytime to night-time concentration ratios of  
12 isoprene did not exhibited the strong light dependency typically associated with atmospheric  
13 isoprene concentrations, which could be attributed to the characteristics of sources of these  
14 species that are discussed in subsequent sections. The temperature and photoactive radiation  
15 (PAR) measurements at Welgegund was used in the MEGAN BVOC emission model, which  
16 indicated that the measurement time (11:00 to 13:00 local time) captured most of the period of  
17 maximum isoprene emission (typically about 12:00 to 2:00 local time). In addition, by  
18 assuming a typical diurnal variation in VOC oxidation rate and boundary layer height, it was  
19 also found that the isoprene concentration of the measurement time is representative of the  
20 daytime isoprene concentration (Greenberg et al., 1999). In Table 3, the concentrations of  
21 BVOC species measured during other campaigns in South Africa and the rest of the world are  
22 presented.

23

24 **Insert Table 2**

25

26 **Insert Table 3**

27

28 The most abundant species observed throughout the study was the monoterpene,  $\alpha$ -pinene, and  
29 the total monoterpene concentration was at least an order of magnitude higher compared to the  
30 concentrations of other BVOC categories. The total annual median (IQR) monoterpene  
31 concentration was 120 (73-242) pptv during the first campaign and 83 (54-145) pptv during the

1 second campaign. As indicated in Table 2,  $\alpha$ -pinene, p-cymene and limonene were the  
2 predominant compounds measured during the first campaign, constituting more than 63 % of  
3 the ambient monoterpene concentrations, while during the second campaign, the dominant  
4 monoterpenes were  $\alpha$ -pinene, limonene and terpinolene, constituting more than 70 % of the  
5 ambient monoterpene concentrations. BVOC flux measurements conducted by Greenberg et  
6 al. (2003) during SAFARI 2000 at a mopane woodland in Botswana indicated that 60 % of the  
7 monoterpene flux was attributed to  $\alpha$ -pinene, while limonene and  $\beta$ -pinene contributed almost  
8 all of the rest of the monoterpenes. Various studies in other regions have also indicated that  $\alpha$ -  
9 pinene is the dominant monoterpene in ambient air reflecting the ubiquitous nature of its  
10 emission (Hellén et al., 2012b; Hakola et al., 2012; Noe et al., 2012). During the AMMA  
11 experiment, Saxton et al. (2007) also detected several monoterpenes in ambient air at Djougou  
12 with concentrations generally higher than monoterpene concentrations recorded by Serca et al.  
13 (2001) (less than 20 pptv) during EXPRESSO at a forest in Northern Congo. Monoterpene  
14 concentrations reported for boreal forest (Hakola et al., 2009; Hakola et al., 2000; Rinne et al.,  
15 2000; Rinne et al., 2005; Rantala et al., 2015; Räisänen et al., 2009; Eerdekens et al., 2009;  
16 Lappalainen et al., 2009), hemiboreal mixed forest (Noe et al., 2012), temperate (Spirig et al.,  
17 2005; Stroud et al., 2005; Fuentes et al., 2007; Mielke et al., 2010), Mediterranean (Davison et  
18 al., 2009; Harrison et al., 2001) and tropical (Rinne et al., 2002) ecosystems ranged between 40  
19 and 7 200 pptv (Table 3). Therefore, there is a large variation in the monoterpene  
20 concentrations measured in different ecosystems, with concentrations measured at Welgegund  
21 being in the low to mid-range. Unlike isoprene that is approximately 10 times lower than  
22 isoprene levels at other ecosystems in the world, the mean monoterpene concentration at  
23 Welgegund is comparable to the previous studies at other ecosystems summarised in Table 3.

24 The annual median (IQR) isoprene concentration measured during the first campaign was 14 (6-  
25 35) pptv, while the annual median (IQR) isoprene concentration measured during the second  
26 sampling campaign was 14 (7-24) pptv. The highest isoprene concentration, i.e. 202 pptv, was  
27 recorded in summer (wet season). Harley et al. (2003) reported that the maximum isoprene  
28 concentration measured during an eight-day campaign in the wet season at a *Combretum-  
29 Acacia* savannah in southern Africa was 860 pptv with a mean midday concentration of 390  
30 pptv, which is considerably higher than isoprene levels measured at Welgegund. Ambient  
31 BVOC measurements conducted by Saxton et al. (2007) at a rural site near Djougou, Benin in  
32 June 2006 during the AMMA project indicated isoprene concentrations >3 000 pptv. Grant et

1 al. (2008) conducted VOC measurements at a small rural Senegalese village during September  
2 2006 that was also a sampling location for the AMMA project and reported that isoprene, which  
3 had a mean concentration of  $300 \pm 100$  pptv, was the only biogenic hydrocarbon present in all  
4 air samples. Serca et al. (2001) reported ambient the mean isoprene concentration for a tropical  
5 forest of Northern Congo during the EXPRESSO study to be  $1820 \pm 870$  pptv at the beginning  
6 of the wet season and  $730 \pm 480$  pptv at the end of the wet season. Nakashima et al. (2014)  
7 reported that the mean isoprene concentration at the Manitou Experimental Forest (MEF) was  
8  $68 \pm 69$  pptv. In general, mean isoprene concentrations measured at Welgegund were at least  
9 an order of magnitude smaller compared to other isoprene measurements in South Africa,  
10 Africa and most other parts of the world.

11 The annual median (IQR) MBO concentrations measured during the first and second campaign  
12 were 7 (3-16) and 4 (3-10) pptv, respectively. MBO and isoprene are both produced from  
13 dimethylallyl diphosphate (DMADP) (Gray et al., 2011). Guenther (2013) indicated that MBO  
14 is emitted from most isoprene emitting vegetation at an emission rate of  $\sim 1$  % of that of  
15 isoprene. However, MBO measured at Welgegund was approximately 30 % of the isoprene  
16 concentrations, which indicated that the main source of MBO at Welgegund is not from  
17 isoprene emitters, but from other MBO emitters. MBO concentration measurements at Manitou  
18 Experimental Forest, USA were  $1\,346 \pm 777$  pptv (Nakashima et al., 2014), which is three  
19 orders of magnitude higher compared to the MBO levels measured at Welgegund. According  
20 to the knowledge of the authors, there are no previous ambient MBO concentrations measured  
21 for Africa.

22 Most SQTs are highly reactive species and are difficult to detect in ambient air samples, which  
23 resulted in concentrations of these species being frequently below the detection limit of the  
24 analytical procedure. This is also reflected in the concentrations of these species being an order  
25 of magnitude lower compared to the other BVOC species measured in this study. The total  
26 annual median (IQR) SQT concentration measured during the first sampling campaign was 8  
27 (5-14) pptv and 4 (3-11) pptv during the second sampling campaign. The most abundant SQT  
28 during the first sampling campaign was longicyclene with an annual mean concentration of 4  
29 (1-4) pptv. During the second sampling campaign,  $\alpha$ -humulene was the most abundant SQT  
30 with an annual mean concentration of 3 (1-5) pptv.

31 The lower BVOC concentrations measured at Welgegund compared to other regions can mainly  
32 be attributed to the much lower isoprene concentrations measured. However, monoterpenes that

1 are important for SOA formation are similar to levels thereof in other environments. In an  
2 effort to explain the BVOC concentrations measured at Welgegund, a comprehensive  
3 vegetation study was conducted, as described in section 2.2. The influence of the type of  
4 vegetation in the region surrounding Welgegund on ambient BVOC concentrations will be  
5 further explored.

6 Jaars et al. (2014) presented concentrations of aromatic VOCs measured at Welgegund during  
7 the same two sampling campaigns discussed in this paper. The total BVOC concentrations  
8 measured were at least an order of magnitude lower compared to concentrations of aromatic  
9 VOCs measured at Welgegund. The most abundant aromatic compound, toluene, had a median  
10 value of 630 pptv, while the most abundant BVOC measured,  $\alpha$ -pinene, had a median value of  
11 37 pptv. In addition, the median of the concentrations of the all the monoterpene species (120  
12 and 83 pptv) was approximately six times lower compared to toluene concentrations (Jaars et  
13 al., 2014).

### 14 **3.3 Seasonal variations**

15 In Figure 5, the panels on the left show monthly median concentrations of (a) isoprene, (b)  
16 MBO, (c) monoterpenes and (d) SQT measured for the two campaigns, while the panels on the  
17 right present the wet (October to April) and dry (May to September) season concentrations of  
18 the respective compounds measured for the two campaigns. As indicated in Section 3.2,  
19 isoprene measured during the first sampling campaign had higher median (mean) daytime  
20 concentrations compared to median (mean) night-time concentrations, which reflects the light  
21 dependency expected from isoprene. All other BVOCs with the exception of two monoterpenes  
22 and one SQT did not indicate significant differences between daytime and night-time median  
23 (mean) concentrations. Therefore the seasonal plots of only isoprene were separated between  
24 daytime and night-time median concentrations. Seasonal variations in BVOC concentrations  
25 are expected due to the response of emissions to changes in environmental conditions, e.g.  
26 temperature and rainfall, as discussed in section 3.1, and the associated biogenic activity. In  
27 addition, BVOC emission is expected to be lower during the winter months (June to August),  
28 since foliar densities rapidly decrease due to deciduous trees dropping their leaves in winter  
29 (Otter et al., 2002). As expected, it is evident that the concentrations of all the BVOC species,  
30 with the exception of the isoprene (Figure 5a), and SQT values (Figure 5d) measured during  
31 the second sampling campaign, were higher in the wet season. The wet season also had more  
32 occurrences of BVOC concentrations that were higher than the range of the box and whisker

1 plot (whiskers indicating  $\pm 2.7\sigma$  or 99.3 % coverage if the data have a normal distribution). In  
2 an isoprene and monoterpene emissions modelling study for southern Africa conducted by Otter  
3 et al. (2003), it was estimated that BVOC emissions will decrease by as much as 85 % in the  
4 dry winter season for grassland and savanna regions. BVOC concentrations measured in this  
5 study indicated much lower decreases from summer (December to February) to winter (June to  
6 August), with isoprene and monoterpene decreasing by only 37 and 29 %, respectively during  
7 the first sampling campaign, while isoprene and monoterpene decreased by only 42 and 23 %,  
8 respectively during the second sampling campaign. This can partially be attributed to the  
9 significant transformation of this biome, as discussed in section 2.2, with large areas  
10 transformed to cultivated land, as indicated in Figure 2. In addition, the study by Otter et al.  
11 (2003) was conducted for the entire southern African region.

12

### 13 **Insert Figure 5**

14

15 The monthly median isoprene concentrations (Figure 5a) measured during the first sampling  
16 campaign indicated the expected seasonal pattern with higher isoprene concentrations  
17 coinciding with the wet and warmer months, with the exception of April that had lower isoprene  
18 concentrations. Surprisingly, during the second sampling campaign, there was no distinct  
19 seasonal pattern observed. However, higher isoprene concentrations seem to coincide with  
20 higher wind speeds (Figure 3d), which are observed for both sampling campaigns. This  
21 indicates that the major sources of isoprene measured at Welgegund can be considered not to  
22 be within close proximity. However, since oxidation products of isoprene (e.g. methyl vinyl  
23 ketone, methacrolein) were not measured in this study, more distant sources of isoprene could  
24 not be verified. It is evident from Figure 2 that the region in close proximity of Welgegund in  
25 the south-western to north-eastern sector largely comprises cultivated land, while in the north-  
26 eastern to south-western sector the predominant land coverage is grassland and natural  
27 vegetation. It is expected that isoprene emissions from the cultivated land will be lower  
28 compared to savanna grassland (Otter et al., 2003). Therefore, if Welgegund is more frequently  
29 affected by winds from the south-western to north-eastern sector, higher wind speeds will  
30 coincide with higher isoprene levels, since the savanna grassland fetch region is distant from  
31 Welgegund and related to the approximately three-hour atmospheric lifetime of isoprene due to  
32 OH radicals.



1 In Figure 6, the wind roses for the BVOCs species measured in this study are presented. It is  
2 evident that the highest isoprene concentrations for the first sampling period were associated  
3 with winds originating from the south to south-western sector, i.e. predominantly from the  
4 grassland region in close proximity during the first sampling campaign resulting in a relatively  
5 more distinct seasonal pattern for isoprene levels. During the second sampling campaign,  
6 higher isoprene concentrations were associated with winds originating from the south-western  
7 to the northern sector, i.e. from the cultivated land area. Therefore, isoprene concentrations  
8 measured during the second sampling period coincided predominantly with stronger wind  
9 speeds from more distant fetch regions.

10

### 11 **Insert Figure 6**

12

13 Distinct seasonal patterns are observed for MBO (Figure 5b) concentrations during both  
14 sampling campaigns, i.e. higher MBO concentrations coinciding with wet warm months and  
15 lower levels corresponding with dry cold months (Figure 3). The MBO concentrations also  
16 corresponded to the seasonal CO<sub>2</sub> uptake (Figure 4). It is also evident from Figure 5b that MBO  
17 concentrations during the wet season in the first sampling campaign were higher compared to  
18 the second sampling campaign, especially from February to April 2011. As mentioned in  
19 section 3.1, the soil moisture measured at a depth of 20 cm (Figure 3g) during the first sampling  
20 campaign was significantly higher from February to August compared to the second sampling  
21 campaign. Therefore, these increased MBO levels measured during the first sampling  
22 campaign can be attributed to increased emissions from deep-rooted plants, e.g. shrubs and  
23 trees. In addition to decreased biogenic activity in the dry winter, the conversion of MBO to  
24 isoprene in the atmosphere could also lead to decreased MBO levels during this period. Jaoui  
25 et al. (2012) reported that MBO conversion to isoprene increased by an order of magnitude  
26 during dry conditions compared to humid conditions. This can also contribute to elevated  
27 isoprene concentrations measured during the dry months at Welgegund (Figure 5a).

28 No distinct seasonal pattern is observed for monoterpene and SQT concentrations, with the  
29 exception of significantly higher levels measured from February to April 2011 during the first  
30 sampling campaign. These increased monoterpene and SQT concentrations can also be  
31 attributed to the significantly higher soil moisture measured at a depth of 20 cm during the first  
32 sampling campaign (Figure 3g), as observed for the MBO. The monoterpene and SQT

1 concentrations measured during the first sampling campaign were generally higher compared  
2 to the second sampling campaign. In Figure S1 (a) and (b) the relationship between soil  
3 moisture and monoterpene concentrations, as well as between soil moisture and SQT are  
4 presented, respectively. It is evident that higher concentrations of monoterpene and SQT are  
5 associated with higher soil moisture measured at a depth of 5 and 20 cm. Otter et al. (2002)  
6 also reported a more pronounced seasonal pattern for isoprene compared to monoterpene  
7 emissions at the Nylsvley Nature Reserve, which is approximately 200 km north-west from  
8 Welgegund.

### 9 **3.4 BVOC emissions from surrounding vegetation**

10 As discussed in section 2.2 and indicated in Figure 2, Welgegund is situated in a region that has  
11 been significantly transformed through cultivation. Cultivated land within the demarcated 60  
12 km radius (Figure 2) consists mainly of maize and, to a lesser degree, sunflower production.  
13 These cultivated lands are also typically characterised by eucalyptus trees, which have a very  
14 high BVOC emission potential (Kesselmeier and Staudt, 1999), planted on their peripheries as  
15 is evident in Figure 2. The grassland region in close proximity of Welgegund (south-western  
16 to north-eastern sector) has a high diversity of grass and woody species, as mentioned in section  
17 2.2. In general, it can be considered that the woody species in the grasslands are major sources  
18 of all the BVOCs measured in this study. Otter et al. (2003) also considered woody vegetation  
19 to be the most important in terms of BVOC emissions in southern Africa. It is generally  
20 considered that crops and grass have very low isoprene-emitting capacities (Kesselmeier and  
21 Staudt, 1999; Guenther, 2013). However, Schuh et al. (1997) indicate that sunflowers emit  
22 isoprene; the monoterpenes  $\alpha$ -pinene,  $\beta$ -pinene, sabinene, 3-carene and limonene; and the  
23 sesquiterpene  $\beta$ -caryophyllene predominantly. In addition, Chang et al. (2014) (with references  
24 therein) also indicated that isoprene has anthropogenic sources in urban areas, which indicates  
25 that the surrounding towns can also contribute to the isoprene concentrations.

26 In an effort to determine possible sources of BVOC species, concentration roses were  
27 compiled, as presented in Figure 6. In general, the concentration roses indicated that isoprene  
28 concentrations were higher from the western sector (indicated by the average and highest  
29 concentrations) that is considered to be a relatively clean regional background region with no  
30 large anthropogenic point sources (Figure 1), while wind direction did not indicate any  
31 significant differences in the concentrations of the other BVOC species. On occasion, higher

1 MBO, monoterpene and SQT concentrations were observed from the south-eastern region,  
2 which may be attributed to a large eucalyptus plantation approximately 15 km south-east from  
3 Welgegund, indicated in Figure 2. However, high isoprene emissions are also usually associated  
4 with eucalyptus trees, which are not observed in the isoprene concentration roses. Therefore,  
5 other sources of MBO, monoterpene and SQT in these regions are most likely to be the main  
6 sources, which can possibly include the urban footprint indicated in this region. In addition,  
7 pine trees are common foreign tree species that are planted on farms in this region (Rouget,  
8 2002), which could be potential sources of MBO and monoterpenes.

9 The similar concentration roses determined for monoterpenes and SQT during the first sampling  
10 campaign can be attributed to similar sources of these species. However, most SQTs have short  
11 atmospheric lifetimes ( $< 4$  min) (Atkinson and Arey, 2003a), which indicated similar sources  
12 within close proximity ( $\sim 1 - 2$  km radius) of Welgegund. Gouinguéné and Turlings (2002)  
13 indicated the emissions of several SQT from young maize plants by testing the effects of soil  
14 humidity, air humidity, temperature, light and fertilisation rate on the emission of BVOCs from  
15 these plants. Therefore, maize production may be a source of monoterpenes and SQT. The  
16 higher SQT concentrations in the south-west and north-west can most likely be attributed to  
17 smaller eucalyptus plantations within a 1 to 2 km radius, as indicated in Figure 2. The high  
18 monoterpene concentrations determined during the second sampling campaign may be  
19 associated with specific monoterpene emitting plants in the region.

20 Floral emissions could also be considered a potential source of monoterpenes in this region,  
21 which could also contribute to the relatively abundance of monoterpenes compared to the  
22 relatively low isoprene concentrations. Floral emissions in this region would typically occur  
23 with the onset of the wet season in October up until February. It is well-known that meadows,  
24 i.e. grazed grasslands in South Africa/this region have a significant number of species that  
25 flower. South African grasslands are considered to be exceptionally species rich (Siebert,  
26 2011), since it is ancient, primary grasslands, i.e. not man-made (Bond, 2016).

27 Of particular interest is the potential sources of 4-allylanisole (estragole) due to its relatively  
28 substantial levels as indicated in Table 2. Bouvier-Brown et al. (2009a) and Misztal, et al.,  
29 (2010) indicated that this species could potentially have a significant contribution to regional  
30 atmospheric chemistry due to relatively large estragole emissions measured from ponderosa  
31 pine trees and oil palms, respectively. As mentioned previously, pine trees are typically found  
32 on farms in this region as intruder tree species (Rouget, 2002), while numerous palm trees

1 occurs in cities/towns surrounding Welgegund (Lubbe *et al.*, 2011). In addition, *Foeniculum*  
2 *vulgare* (fennel) – considered a typical source of estragole – is an abundant and common weed  
3 in this study region (Lubbe *et al.*, 2010). Furthermore, estragole emissions could also potentially  
4 have a floral origin.

5 Although a comprehensive vegetation survey has been conducted within a 60 km radius of  
6 Welgegund, vegetation types have been identified only generally at this stage, as indicated in  
7 section 2.2. Therefore, the predominant woody species in each of the regions surrounding  
8 Welgegund associated with specific BVOC emissions have not yet been characterised.

### 9 **3.5 Statistical correlations**

10 Spearman's correlation analyses were applied to correlate the measured concentrations of  
11 isoprene, MBO, monoterpenes and SQT measured to each other in order to substantiate sources  
12 of these species. These correlations for the two sampling campaigns are presented in Table 4,  
13 with correlations in the wet seasons listed in the lower bottom (not highlighted) and correlations  
14 in the dry season presented in the top right (bold highlighted). It is evident that MBO had good  
15 correlations with monoterpenes and SQT in the wet season, as well as with monoterpenes in  
16 the dry season during the first sampling campaign. Although not as distinct as during the first  
17 sampling campaign, MBO did also correlate with monoterpenes during the wet and dry season,  
18 as well as with SQT in the dry season during the second sampling campaign. During the first  
19 sampling campaign, monoterpenes had a strong correlation with SQT in the wet season and  
20 moderate correlation during the dry season, while strong correlations between monoterpene and  
21 SQT were determined in the dry season and a moderate correlation during the wet season during  
22 the second sampling campaign. As indicated previously, concentration roses did indicate  
23 similar sources of MT and SQT, especially during the first sampling campaign, which is  
24 signified by these correlations.

25

#### 26 **Insert Table 4**

27

28 Spearman correlations between BVOCs and other parameters measured at Welgegund did not  
29 show significant correlations. However, in certain instances, good correlations were observed  
30 between soil moisture and MBO, monoterpenes and SQT concentrations. This is expected,

1 since the monthly average concentrations of these species indicated increased levels thereof  
2 that were associated with increased soil moisture from February to April 2011. Therefore, in  
3 an effort to further statistically explore the dataset, explorative multilinear regression was  
4 performed by using all ancillary measurements as input data in order to indicate parameter  
5 interdependencies on the BVOC concentrations measured. In Figure 7, the root mean square  
6 error (RMSE) difference between the calculated and measured BVOC concentrations, as a  
7 function of the number of independent variables included in the optimum MLR solution, is  
8 presented. It is evident that interdependence between temperature, soil temperatures and PAR  
9 yielded the largest decrease in RMSE for isoprene concentrations measured. However, for  
10 MBO, monoterpenes and SQT, a much more significant contribution from soil moisture is  
11 observed to decrease the RMSE differences between calculated and measured BVOC levels. It  
12 is also evident that the interdependence between soil moisture and soil temperature at 20 cm is  
13 important to estimate MBO, monoterpene and SQT concentrations. Therefore, explorative  
14 MLR indicated that temperature had the largest influence on isoprene concentrations, while soil  
15 moisture was the most significant for MBO, monoterpenes and SQT levels.

16

17 **Insert Figure 7**

18

### 19 **3.6 Reactivity of BVOCs**

20 It is important to evaluate the significance of BVOCs on their atmospheric reactivity, since  
21 these species are important precursor species in the photochemical formation of tropospheric  
22 O<sub>3</sub> and SOA. This is particularly relevant for South Africa, with various recent studies  
23 indicating that O<sub>3</sub> is currently the most problematic pollutant in South Africa (Laakso et al.,  
24 2013; Venter et al., 2012; Beukes et al., 2013). In addition, Vakkari et al. (2015) also indicated  
25 the importance of VOCs for new particle formation and growth. Therefore, the O<sub>3</sub> formation  
26 potential (OFP), reaction rates with O<sub>3</sub> and OH reactivities of the BVOCs measured in this  
27 study were evaluated.

28 The OFP of BVOCs was determined by calculating the product of the average concentration  
29 and the maximum incremental reactivity (MIR) coefficient of each compound, i.e.  $OFP =$   
30  $VOC \times MIR$  (Carter, 2009). The MIR scale has been used to assess OFP for aromatic

1 hydrocarbons in numerous previous studies (Hoque et al., 2008; Jaars et al., 2014; Na et al.,  
2 2005). The reaction rates for reactions between O<sub>3</sub> and BVOCs were calculated with Eq. (3):

$$3 \text{ reaction rates} = k_{X,O_3} [X][O_3], \quad (3)$$

4 where [X] is the BVOC concentration, [O<sub>3</sub>] the ozone concentration and  $k_{X,O_3}$  the reaction rate  
5 constant for the reaction between X and O<sub>3</sub>. Since direct OH reactivity measurements were not  
6 available, the OH reactivities ( $s^{-1}$ ) of the BVOCs were calculated, using Eq. (4):

$$7 \text{ OH reactivity} = k_{X,OH} [X] \quad (4)$$

8 where [X] is the BVOC concentration and  $k_{X,OH}$  the reaction rate constant of the reaction  
9 between X and OH. In Table 5, the OFP calculated for each of the BVOCs measured in this  
10 study, as well as the reaction rate constants for the reactions of these species with O<sub>3</sub> and OH,  
11 are listed.

12

### 13 **Insert Table 5**

14

15 Table 5 indicates that, according to the OFP calculated with MIR coefficients,  $\alpha$ -pinene,  
16 isoprene and p-cymene had the highest OFP in descending order during the first sampling  
17 campaign. During the second sampling campaign,  $\alpha$ -pinene also had the highest OFP, while  
18 limonene and isoprene had the second and third highest OFPs, respectively. A comparison of  
19 the OFP calculated in this study to the OFP calculated by Jaars et al. (2014) for anthropogenic  
20 aromatic hydrocarbons measured at Welgegund (with MIR coefficients) indicates that the OFP  
21 of BVOCs is an order of magnitude smaller than the OFP of aromatic hydrocarbons at  
22 Welgegund. The combined O<sub>3</sub> formation potentials of all the BVOCs measured calculated with  
23 MIR coefficients during the first and second campaign were 1162 and 1022 pptv, respectively.

24 In Figure 8 (a), the monthly mean reaction rates for the reactions between O<sub>3</sub> and BVOCs  
25 measured in this study are presented. Higher reaction rates between BVOCs and O<sub>3</sub> contribute  
26 to increased atmospheric O<sub>3</sub> depletion. Significantly higher reaction rates were calculated for  
27 February 2015. It is evident from Figure 8(a) that  $\alpha$ -pinene and limonene had the highest  
28 reaction rates with O<sub>3</sub>, while isoprene exhibited relatively small contributions to the O<sub>3</sub> depletion.  
29 The other BVOCs also had relatively low reaction rates for their reactions with O<sub>3</sub>. In Figure  
30 8(b), the relative monthly contributions of each of the BVOCs to the total OH-reactivity of

1 BVOCs are presented. It is evident that largest contributions to the OH-reactivity of BVOCs  
2 measured at Welgegund are from limonene,  $\alpha$ -pinene and terpinolene for all of the months  
3 during both sampling campaigns. This is expected, since monoterpenes had the highest  
4 atmospheric concentrations compared to the other BVOCs measured in this study. It is also  
5 evident, especially during the first sampling campaign, that isoprene levels increased with the  
6 onset of spring in September.

7

8 **Insert Figure 8**

9

## 10 **4 Conclusions**

11 The annual median concentrations of isoprene, MBO, monoterpene and SQT determined during  
12 two sampling campaigns indicated that the sum of the concentrations of the monoterpenes was  
13 an order of magnitude higher than the concentrations of other BVOC species, with  $\alpha$ -pinene  
14 being the most abundant species. Although monoterpene concentrations were similar to levels  
15 measured at other regions in the world and in a South Africa, very low isoprene concentrations  
16 at Welgegund led to a significantly lower total BVOC concentration compared to levels  
17 reported in most previous studies. In addition, total BVOC concentrations were an order of  
18 magnitude lower compared to the total aromatic VOC concentrations measured by Jaars et al.  
19 (2014) at Welgegund.

20 Distinct seasonal patterns were observed for MBO during both sampling campaigns, which  
21 coincided with wet and warmer months. Although less pronounced, a similar seasonal trend  
22 than for MBO was observed for isoprene during the first sampling campaign, while higher  
23 isoprene concentrations during the second sampling campaign were associated with higher  
24 wind speeds that indicated a distant source region of isoprene. No distinct seasonal pattern was  
25 observed for monoterpene and SQT concentrations. However, significantly higher levels of  
26 monoterpene and SQT, as well as MBO were measured from February to April 2011 during the  
27 first sampling campaign, which were attributed to the considerably higher soil moisture  
28 measured at a depth of 20 cm resulting for the wet season preceding the first campaign and is  
29 indicative of biogenic emissions from deep-rooted plants.

30 Woody species in the grassland region were considered to be the main sources of BVOCs  
31 measured, while sunflower and maize crops were also considered to be potential sources for

1 BVOCs in this region. Multilinear regression analysis indicated that soil moisture had the most  
2 significant impact on atmospheric levels of MBO, monoterpene and SQT concentrations, while  
3 temperature had the greatest influence on isoprene levels.

4 The O<sub>3</sub> formation potentials of the BVOCs measured were an order of magnitude smaller than  
5 that determined for anthropogenic VOCs measured at Welgegund. Isoprene and the  
6 monoterpenes:  $\alpha$ -pinene, p-cymene, limonene and terpinolene, had the largest contribution to  
7 O<sub>3</sub> formation potential.  $\alpha$ -Pinene and limonene had the highest reaction rates with O<sub>3</sub>, while  
8 isoprene exhibited relatively small contributions to the O<sub>3</sub> depletion. Limonene,  $\alpha$ -pinene and  
9 terpinolene had the largest contributions to the OH-reactivity of BVOCs.

10 It is important in future work that a comprehensive study on BVOC emissions from specific  
11 plant species in the area surrounding Welgegund must be performed in order to relate the  
12 emission capacities of vegetation types to the atmospheric BVOCs measured. It is also  
13 recommended that the oxidation products of BVOC species are measured in order to verify  
14 distant source regions of BVOCs measured at Welgegund. In addition, the interactions between  
15 anthropogenic and biogenic VOCs must also be further explored, together with other ancillary  
16 measurements conducted at Welgegund (e.g. SO<sub>2</sub>, NO<sub>2</sub> and O<sub>3</sub>). Future work must also include  
17 investigating the reactions of the measured VOCs with atmospheric oxidants (e.g. •OH and O<sub>3</sub>)  
18 with atmospheric chemistry models in order to establish, for instance, whether O<sub>3</sub> formation  
19 within the region is VOC- or NO<sub>2</sub>-limited.

20

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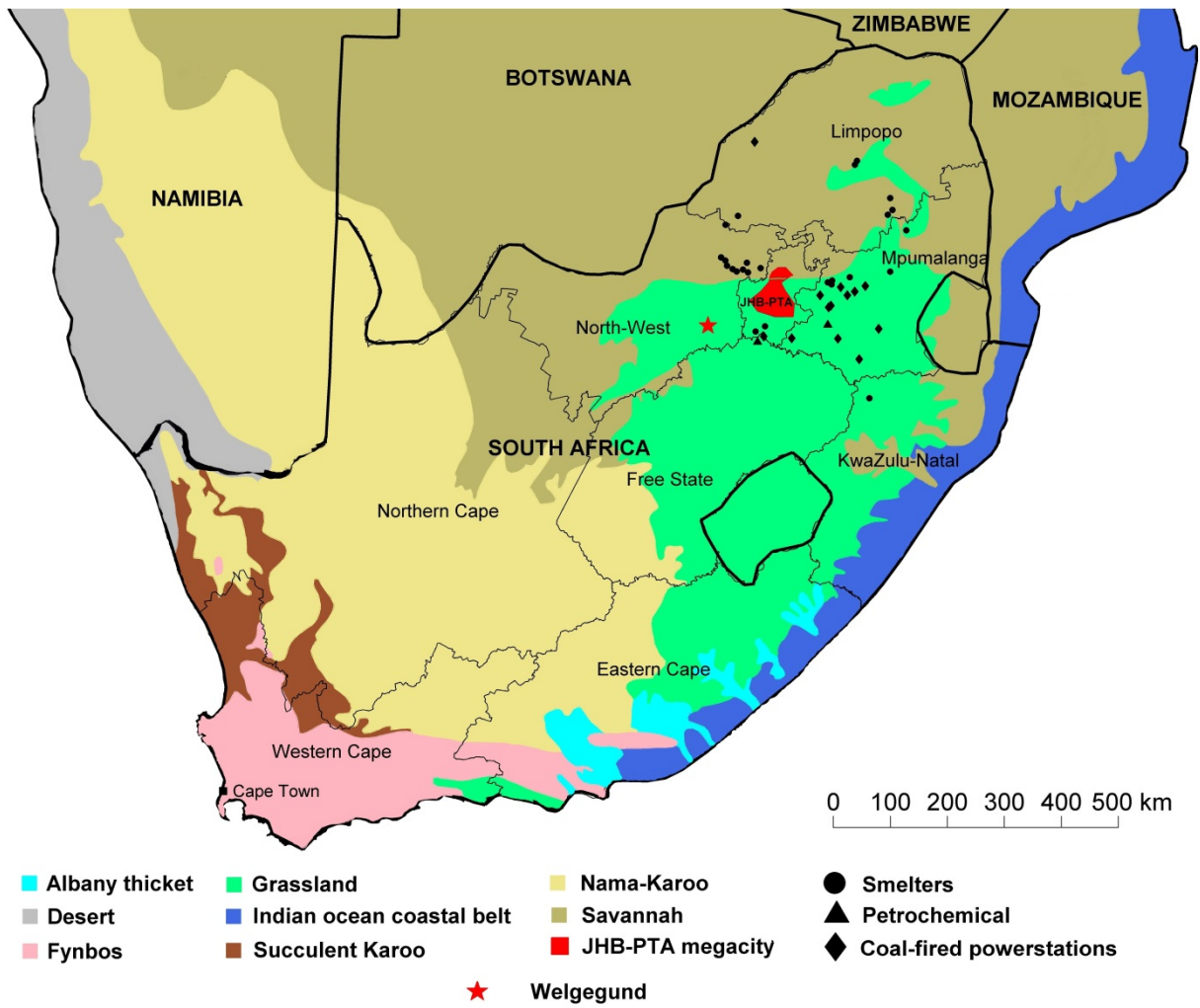
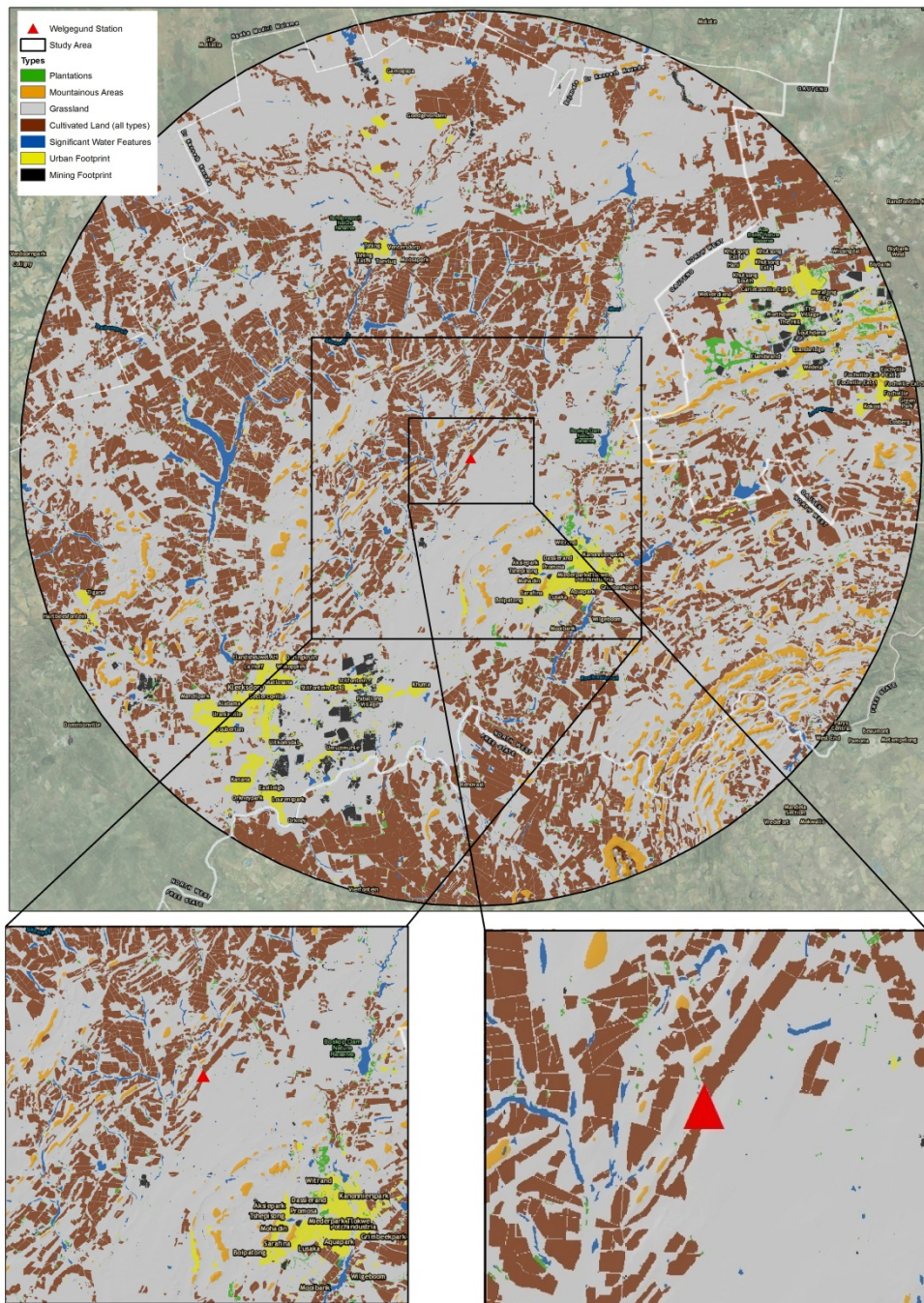


Figure 1. Map of southern African indicating the location of the Welgegund measurement station within the context of bioregion and large point sources in the industrial hub of South Africa (Mucina and Rutherford, 2006).





1

2 Figure 2. General vegetation map for 60 km radius of Welgegund measurement station.

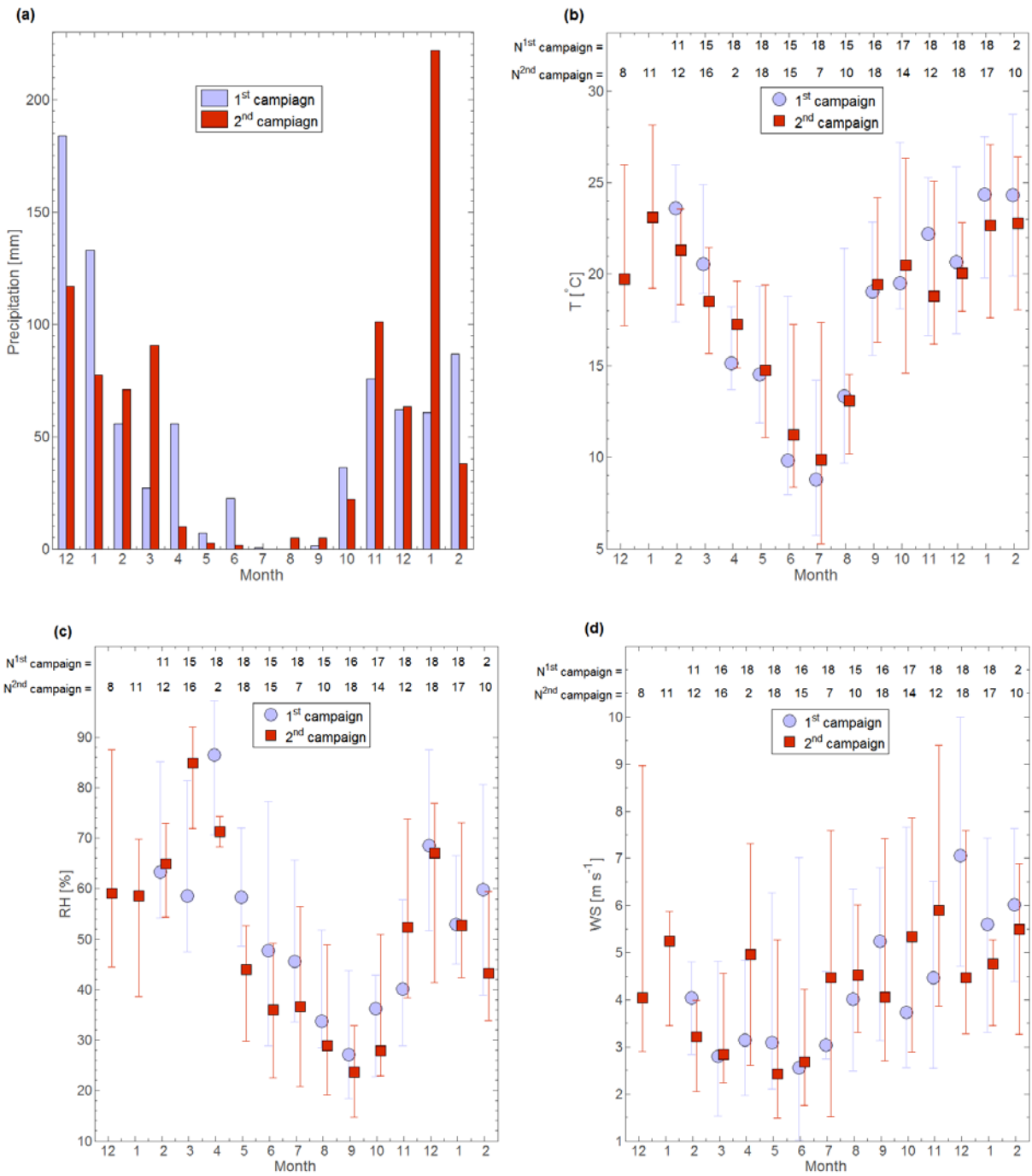
3



1 Table 1. Lifetime ( $\tau$ ) of BVOCs calculated for the average concentration of OH radicals (ca.  
 2 0.04 pptv) as reported by Ciccioli et al. (2014) and the annual average O<sub>3</sub> (ca. 36 ppbv)  
 3 concentration measured for the two campaigns at Welgegund.

		$\tau_{OH}$	$\tau_{O_3}$
	Isoprene	2.8 hr	1 day
	MBO	10.3 hr	7.5 day
Monoterpenes	$\alpha$ -Pinene	5.3 hr	3.6 hr
	Camphene	5.3 hr	14.5 day
	$\beta$ -Pinene	3.6 hr	20.9 hr
	$\Delta^3$ -Carene	3.2 hr	8.5 hr
	p-Cymene	18.8 hr	261.6 day
	1,8-Cineol	12.5 hr	-
	Limonene	1.7 hr	1.6 hr
	Terpinolene	12.6 hr	2.3 hr
	AMCH	2.9 hr	-
	Nopinene	1.4 day	-
	Bornylacetate	1.5 day	-
4-Allylanisole	5.2 hr	1.1 day	
Sesquiterpenes	Longicyclene	1.3 day	-
	iso-Longifolene	2.9 hr	1.1 day
	Aromadendrene	4.5 hr	1.1 day
	$\alpha$ -Humulene	1 hr	21.6 min

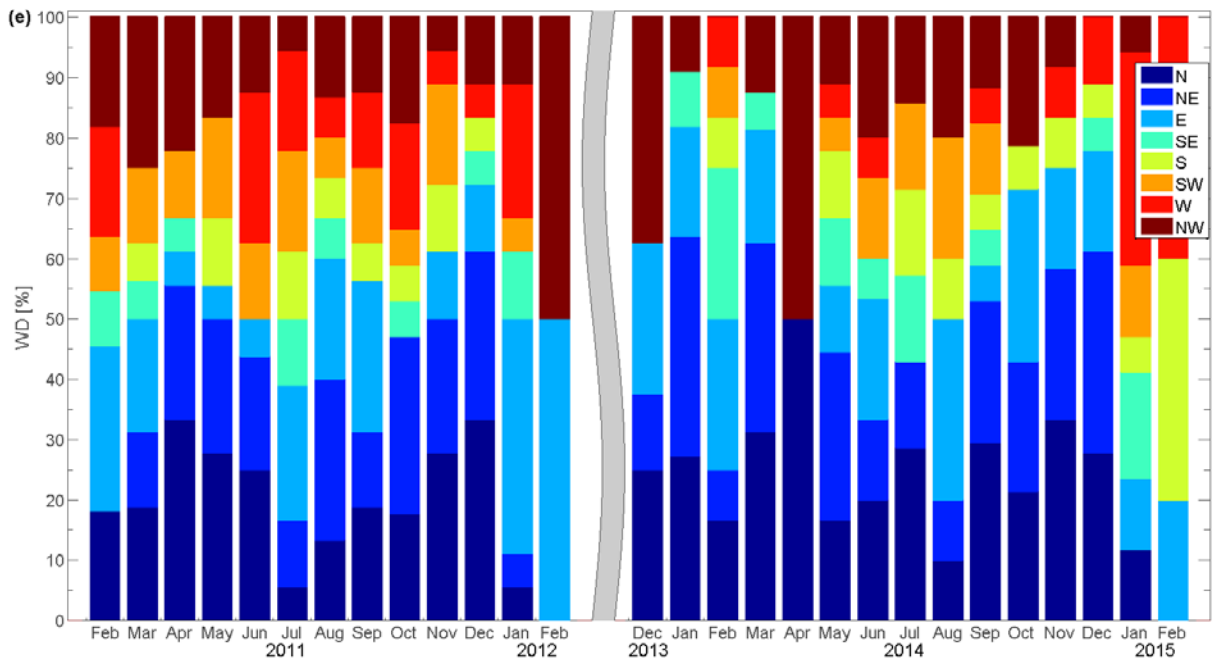
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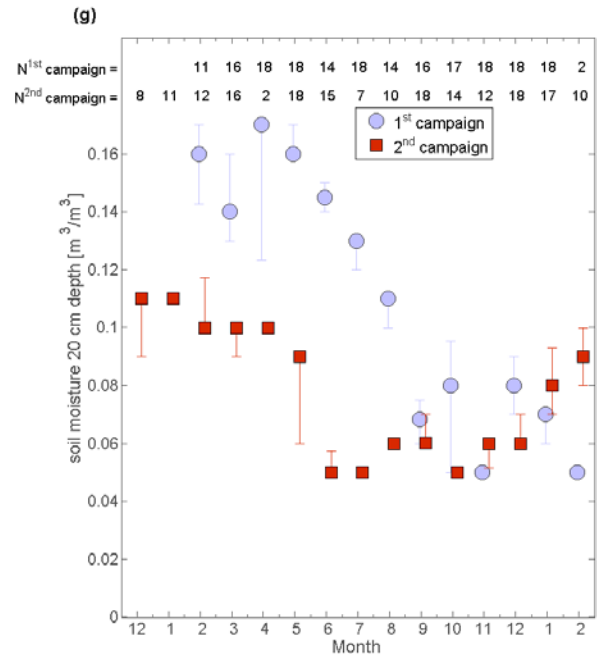
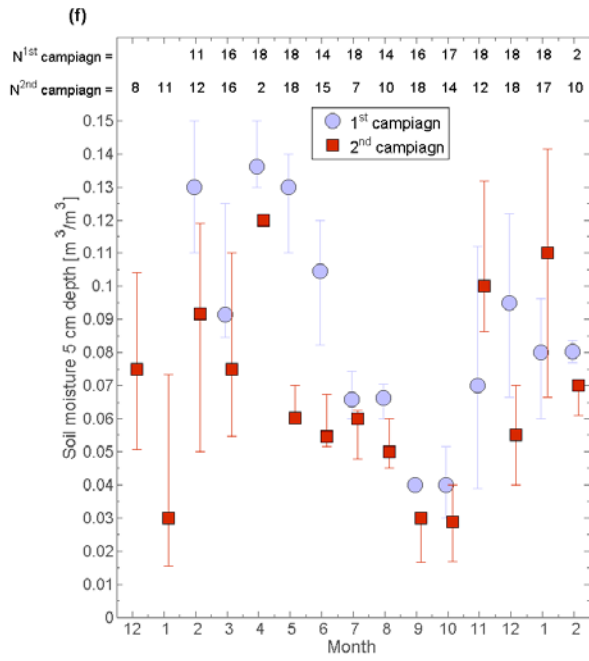
1

2

3 Figure 3. Monthly variation of (a) accumulated precipitation, (b) temperature, (c) relative  
 4 humidity, (d) wind speed, (e) wind direction, and (f) and (g) soil moisture at 5 and 20 cm depth,  
 5 respectively. Error bars indicate upper and lower quartiles.



1

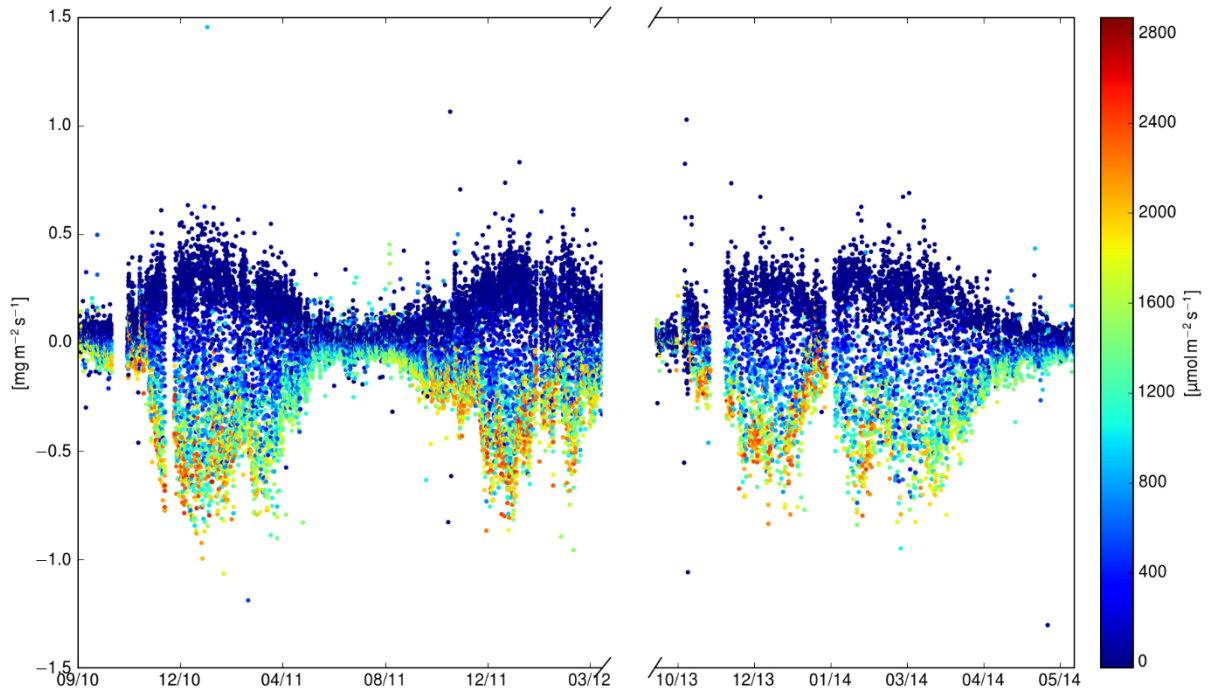


2

3 Figure 3. Continued.

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3 Figure 4. Micrometeorological CO<sub>2</sub> flux measurements at Welgegund (Räsänen et al., 2016).

4 The colour bar indicates the Photosynthetic Photon Flux Density (PPFD).

5

1 Table 2. The ambient BVOC concentration for the two campaigns measured at Welgegend.

2

	First campaign				Second campaign			
	Median (Mean)	IQR (25th - 75th)	Median (Mean) day/night ratio	N	Median (Mean)	IQR (25th - 75th)	Median (Mean) day/night ratio	N
Isoprene	14 (28)	6-35	2.01 (1.7)	187	14 (23)	7-24	0.99 (1.31)	175
MBO	7 (12)	3-16	0.94 (0.91)	178	4 (8)	3-10	1.13 (1.20)	163
<b>Monoterpenes</b>								
$\alpha$ -Pinene	37 (71)	28-83	1.14 (0.93)	197	15 (57)	9-23	1.23 (1.09)	191
Camphene	4 (8)	2-9	1.26 (1.08)	178	2 (4)	1-3	1.20 (0.88)	113
$\beta$ -Pinene	9 (19)	5-18	1.11 (0.98)	195	3 (5)	2-6	1.31 (1.23)	171
$\Delta^3$ -Carene	3 (6)	2-5	1.52 (1.13)	156	2 (4)	1-4	1.13 (0.71)	58
1,8-Cineol	3 (13)	1-7	1.04 (0.92)	162	1 (2)	1-2	0.94 (0.77)	75
Limonene	21 (30)	9-40	1.24 (1.04)	197	16 (54)	8-36	1.23 (0.96)	187
Terpinolene	4 (14)	3-11	1.35 (1.02)	141	22 (28)	16-34	1.57 (1.60)	25
Nopinene	6 (7)	4-9	1.13 (1.09)	167	8 (11)	6-13	1.31 (1.26)	176
Bornylacetate	1 (2)	1-2	1.19 (1.08)	49	2 (3)	1-3	1.40 (1.82)	101
<b>Monoterpene-related BVOCs</b>								
p-Cymene	20 (48)	12-33	1.08 (0.96)	197	7 (15)	4-13	1.20 (0.97)	186
4-Allylanisole	8 (11)	5-13	1.26 (0.96)	118	1 (12)	1-3	1.32 (0.59)	70
AMCH	5 (7)	1-12	0.28 (0.64)	41	3 (4)	2-5	1.33 (1.29)	24
$\Sigma$ monoterpenes and monoterpene-related BVOCs	120 (235)	73-242			83 (198)	54-145		
<b>Sesquiterpenes</b>								
Longicyclene	2 (4)	1-4	1.32 (1.19)	152	1 (2)	1-3	0.95 (0.65)	34
iso-Longifolene	2 (3)	1-4	1.06 (0.89)	52	1 (1)	1	1.19 (1.39)	7
Aromadendrene	1 (1)	1		2	2 (2)	1-3	1.65 (1.91)	73
$\alpha$ -Humulene	1 (1)	1		3	1 (3)	1-5	0.86 (3.53)	4
Alloaromadendrene	2 (3)	1-4	0.96 (0.84)	31				
$\Sigma$ Sesquiterpenes	8 (12)	5-14			4 (8)	3-11		

3

4

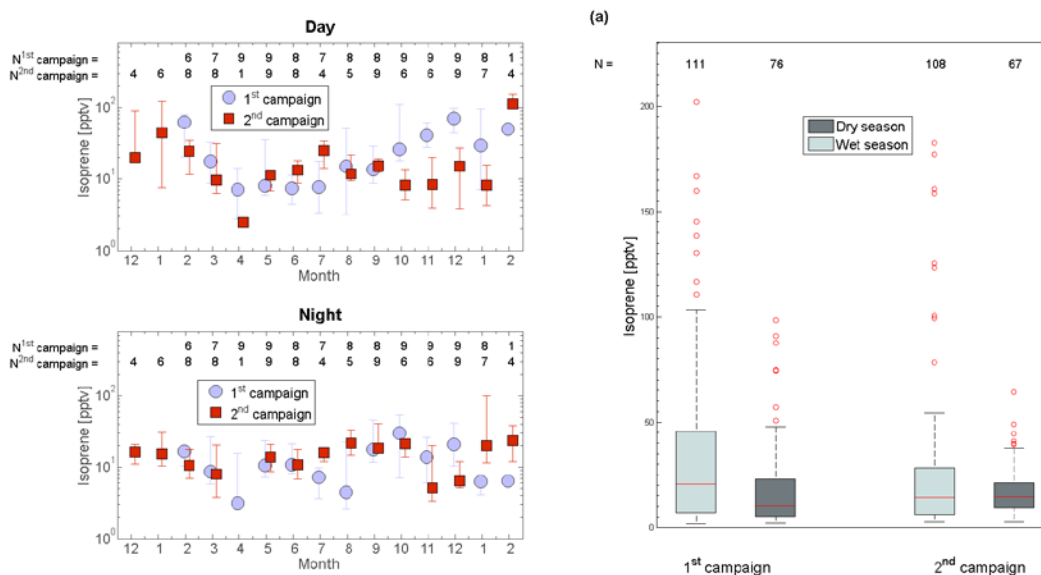
1 Table 3. Ambient BVOC concentrations (pptv) as reported by Noe et al. (2012) for various  
 2 ecosystems and then modified. avg = mean value, med = median value, max = maximal value  
 3 of the measurements reported.

Location	Isoprene	MBO	Monoterpenes	Date	References
<b>Grassland</b>					
Welgegund, SA	28 (avg), 202(max)	12 (avg), 61(max)	235(avg), 1744(max)	Feb 2011-Feb 2012	this study
	23(avg), 182(max)	7 (avg), 47(max)	198(avg), 3081(max)	Dec 2013-Feb 2015	this study
<b>Savannah</b>					
KNP, SA	390(avg),860(max)			Feb 2001	Harley et al. (2003)
Benin	>3000(max)		>5000(max)	7-13 Jun 2006	Saxton et al. (2007)
Village, Senegal	300(avg)			Sept. 2006	Grant et al. (2008)
<b>Boreal</b>					
Hyttiälä, Finland			900(avg), 1800(max)	2000-2007	Hakola et al. (2009)
	40-110		100-700	Apr 2005	Eerdekens et al. (2009)
	220(med),360(max)		300(med), 600(max)	Summer 2006/2007	Lappalainen et al. (2009)
	70(med), 110(max)		200(med), 300(max)	Winter 2006/2007	
	110(avg), 430(max)		100(avg), 2700(max)	Jul 2004	Rinne et al. (2005)
			40-450	37 m, Aug 1998	Rinne et al. (2000)
			140-500	19.5 m, Aug 1998	
			450-630	2 m, Aug 1998	
Huhus, Finland			900(avg), 2160(max)	Jun.-Sep 2003	Räisänen et al., (2009)
Pötsönvaara, Finland	320-1690		1700-3200	Apr.-Oct 1997, 1998	Hakola et al. (2000)
<b>Hemiboreal</b>					
Järvselja, Estonia	360-2520		1800-7200	Spring and Summer 2010	Noe et al. (2012)
	120-200 (med)		400-1400 (med)	Oct 2009-Sep 2010	Noe et al. (2012)
<b>Temperate</b>					
Michigan, USA	2520(avg), 8160(max)		310(avg), 1100(max)	Summer 2008	Mielke et al. (2010)
Jülich, Germany	1980(avg), 10790(max)		250(avg), 1470(max)	Jul 2003	Spirig et al. (2005)
Duke Forest, USA	1500-2200		310-790	Jul 2003	Stroud et al. (2005)
Oak Ridge, USA	5000-15000		500-1600	Jul 1999	Fuentes et al. (2007)
MEF, USA	70(avg)	1346(avg)	0.497(avg)	22-28 Aug. 2008	Nakashima et al. (2014)
<b>Mediterranean</b>					
Castelpoziano, Italy	141-250		100-200	May-Jun 2007	Davison et al. (2009)
AM, Greece	1500(avg), 7900(max)		900(avg), 5000(max)	Jul-Aug 1997	Harrison et al. (2001)

Tropical				
FNT, Brazil	2000(avg), 4000(max)	50(avg), 130(max)	Jul 2000	Rinne et al. (2002)
NNNP, NC	1820±870		16–24 Ma. 1996	Serca et al. (2001)
	730±480		21 Nov–11 Dec 1996	

- 1 SA = South Africa
- 2 WA = West Africa
- 3 KNP = Kruger National Park
- 4 MEF = Manitou Experimental Forest
- 5 AM = Agrafa Mountains
- 6 FNT = Floresta Nacional do Tapajos
- 7 NNNP = Nouabale-Ndoki National Park
- 8 NC = Northern Congo
- 9

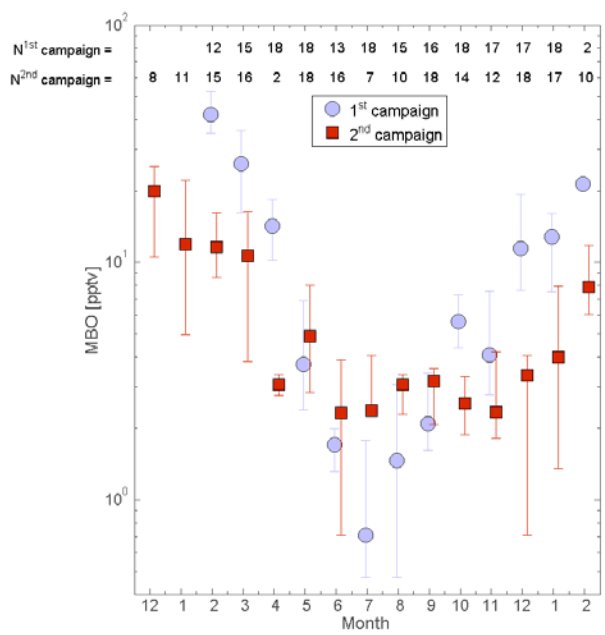
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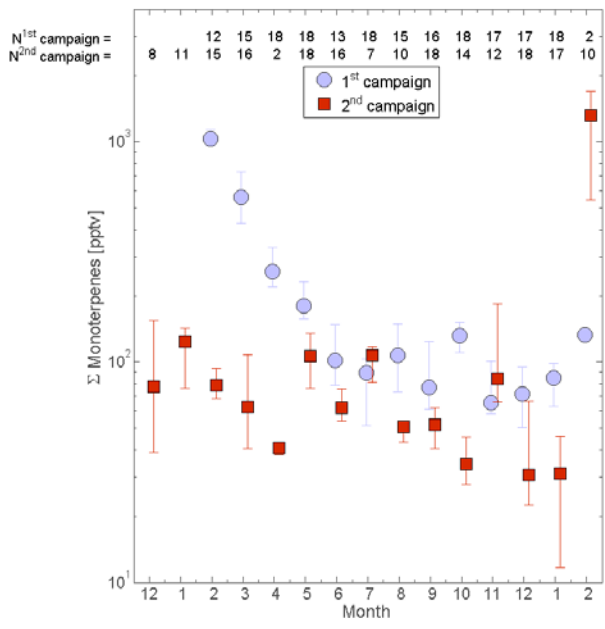
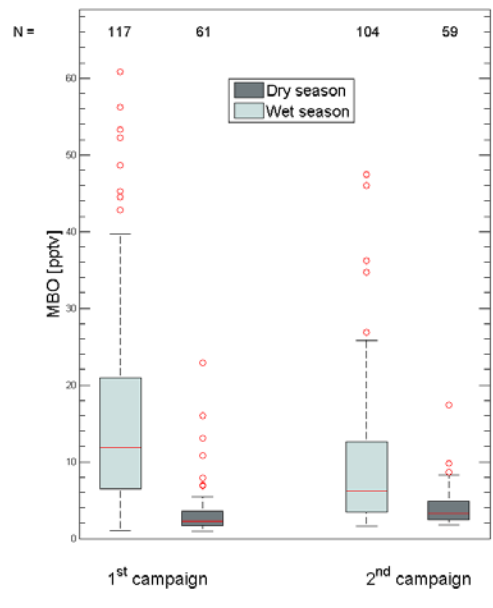
3 Figure 5. The panels on the left show monthly median concentrations of (a) isoprene, (b) MBO,  
4 (c) monoterpenes and (d) SQT measured for the two campaigns. Error bars indicate upper and  
5 lower quartiles. The values displayed near the top of the graphs indicate the number of samples  
6 (N<sup>1st</sup> and N<sup>2nd</sup> campaign) analysed for each month. The panels on the right show the wet and  
7 dry season concentrations of the respective compounds measured for the two campaigns. The  
8 red line of each box indicates the median (50<sup>th</sup> percentile), the black dot the mean, the top and  
9 bottom edges of the box the 25<sup>th</sup> and 75<sup>th</sup> percentiles, the whiskers  $\pm 2.7\sigma$  or 99.3 % coverage if  
10 the data have a normal distribution and the red circles outliers of the range of the box and  
11 whisker plot. The values displayed near the top of the graphs indicate the number of samples  
12 (N) analysed for the wet and dry season.





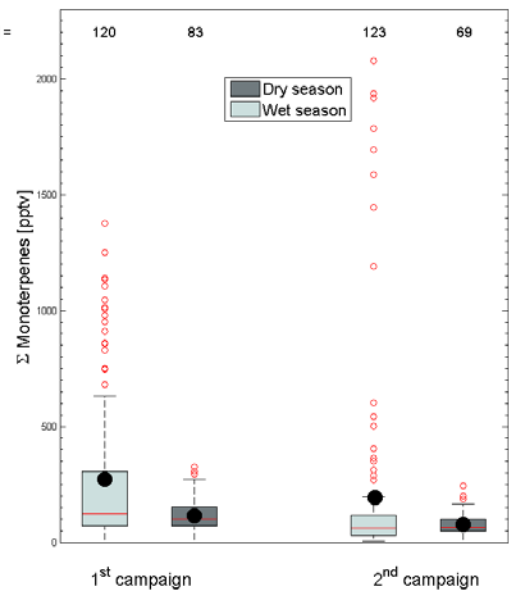
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(b)

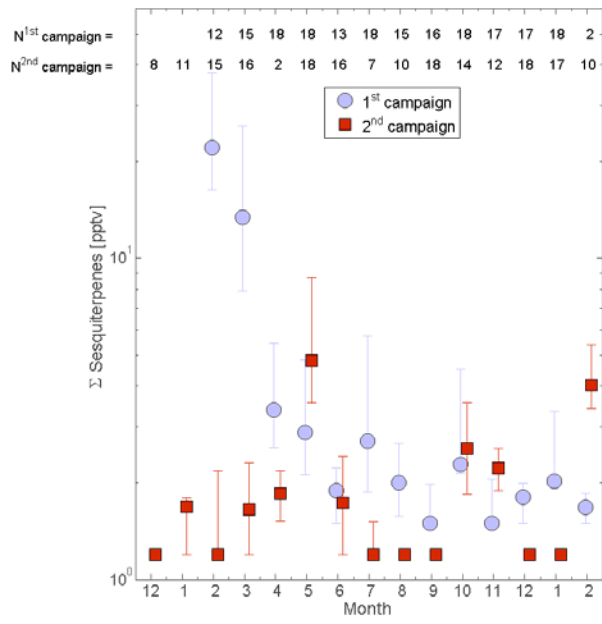


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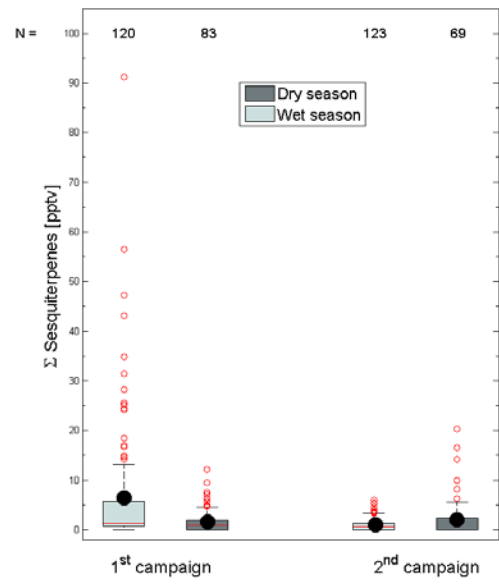
(c)



3 Figure 5. Continued.



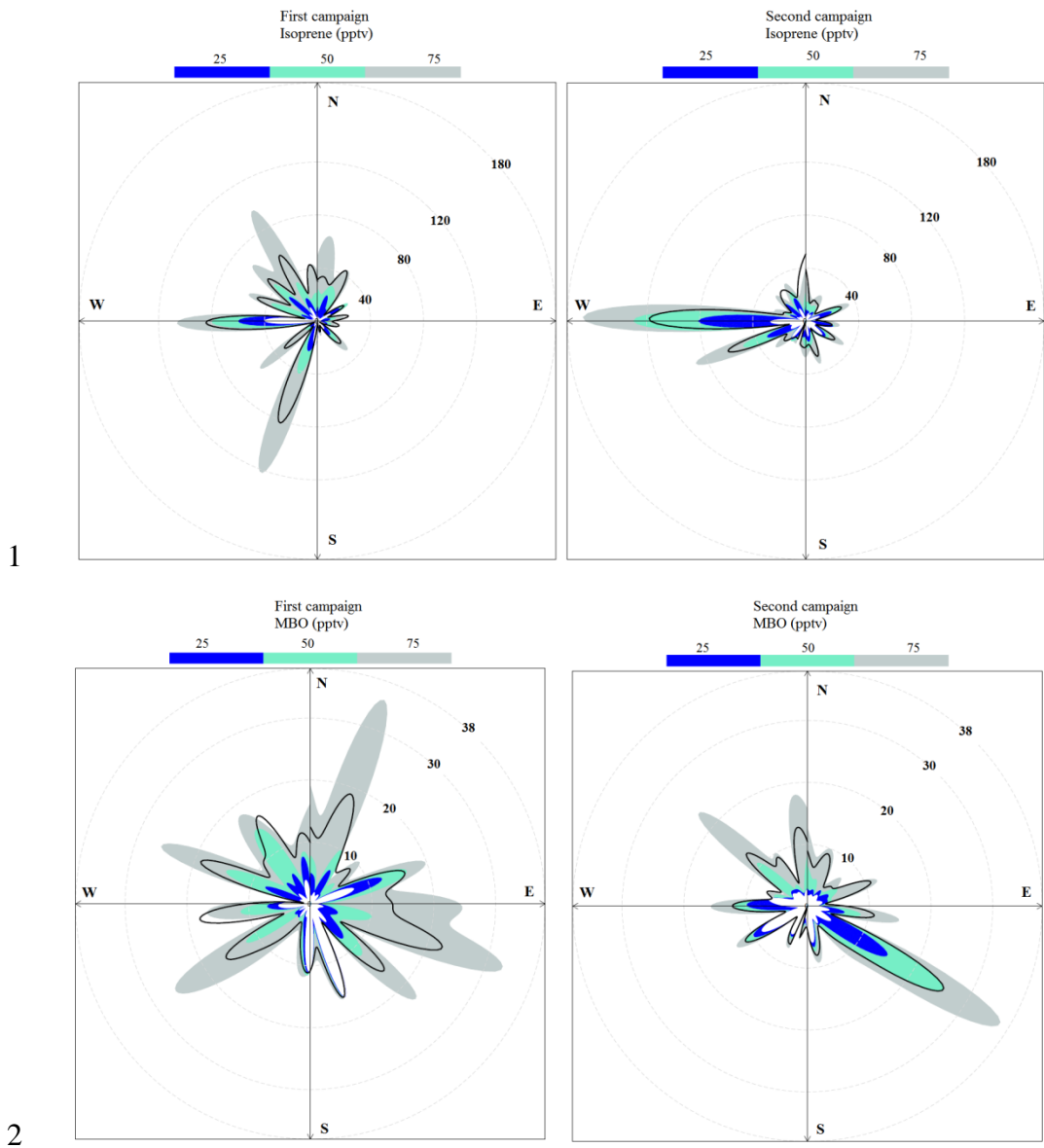
(d)



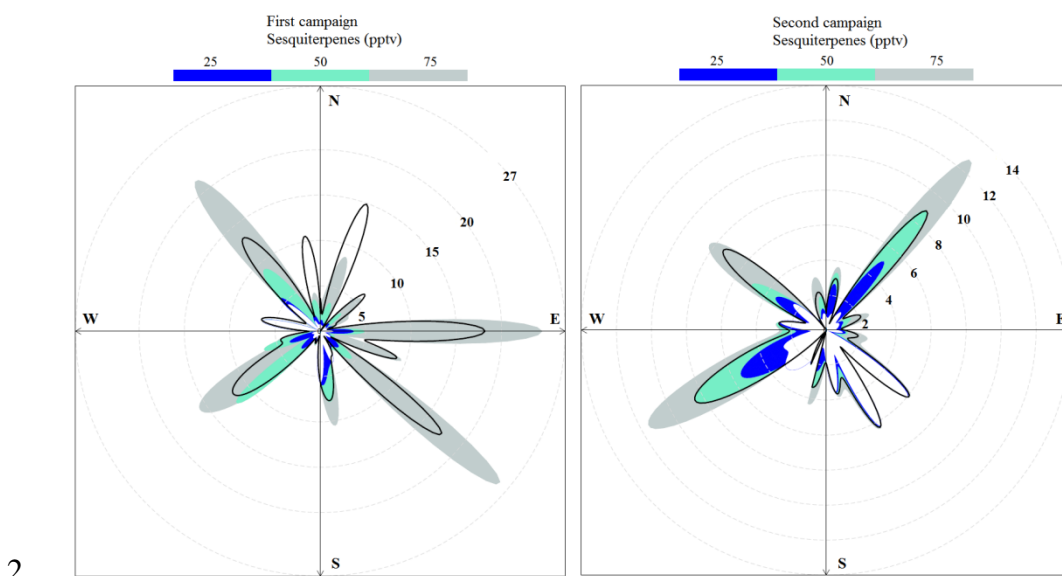
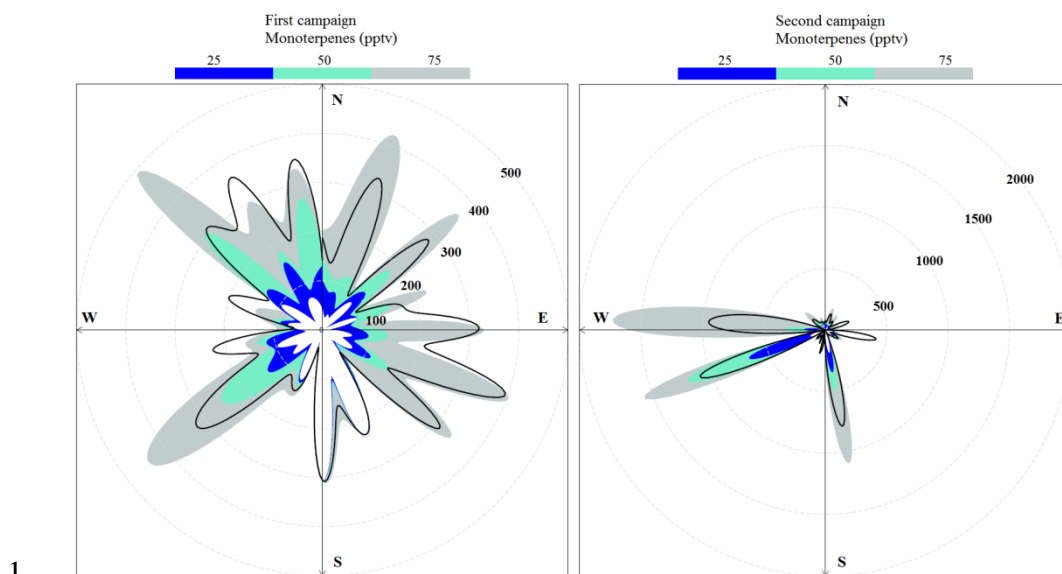
1

2 Figure 5. Continued.

3



3 Figure 6. BVOC concentration rose at Welgegund for the two sampling campaigns. Different  
 4 colours represent percentiles: blue 25 %, aquamarine 50 %, azure 75 % and the black solid line  
 5 the average.



3 Figure 6. Continued.

4

1 Table 4. Spearman's correlation coefficients between the BVOCs during the wet and dry  
 2 season of the first campaign (a) and second campaign (b).

3  
 4  
 5

(a)

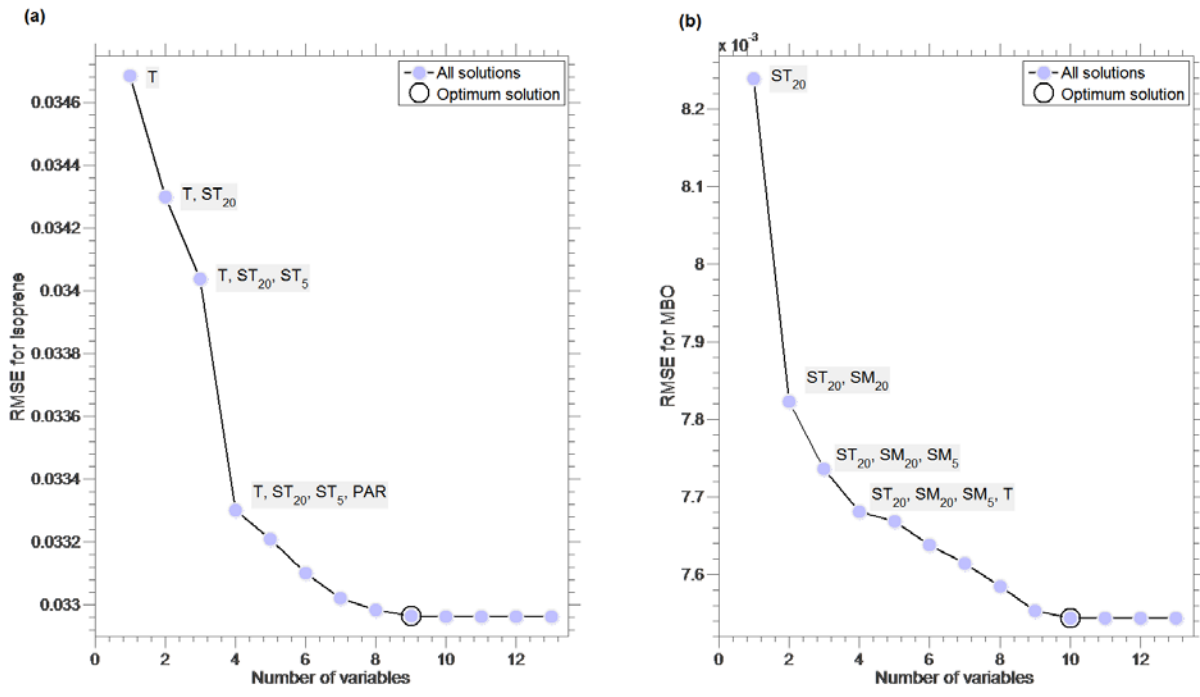
		Dry season			
		Isoprene	MBO	MT	SQT
Wet season	Isoprene	-	<b>0.52</b>	<b>0.03</b>	<b>-0.10</b>
	MBO	0.09	-	<b>0.57</b>	<b>-0.10</b>
	MT	-0.20	0.68	-	<b>0.27</b>
	SQT	-0.04	0.56	0.80	-

6  
 7  
 8

(b)

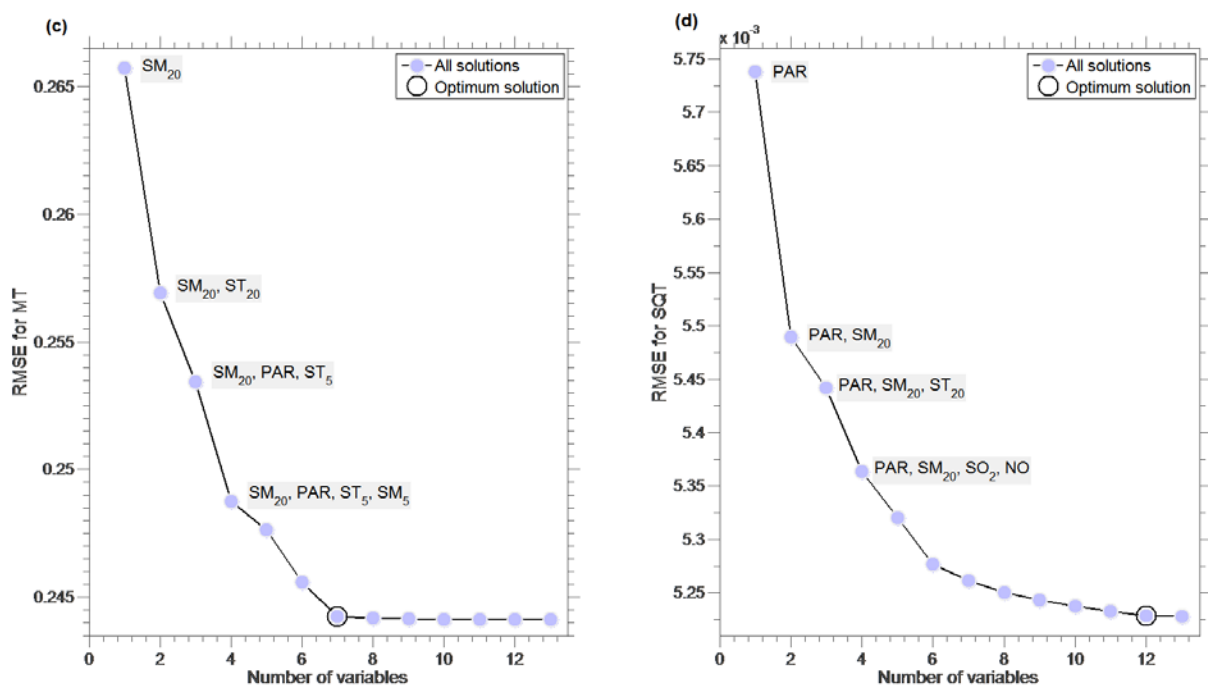
		Dry season			
		Isoprene	MBO	MT	SQT
Wet season	Isoprene	-	<b>0.39</b>	<b>-0.11</b>	<b>0.09</b>
	MBO	0.50	-	<b>0.39</b>	<b>0.48</b>
	MT	0.27	0.38	-	<b>0.60</b>
	SQT	0.20	0.01	0.26	-

9  
 10  
 11



1

2 Figure 7. The optimum combination of independent variables to include in a MLR equation to  
 3 calculate the dependant variable, i.e. BVOC concentrations. The root mean square error  
 4 (RMSE) difference between the calculated and measured concentrations indicated that the  
 5 inclusion of (a) 9 parameters for isoprene, (b) 10 parameters for MBO, (c) 7 parameters for MT,  
 6 and (d) 12 parameters for SQT in the MLR solution was the optimum.



7

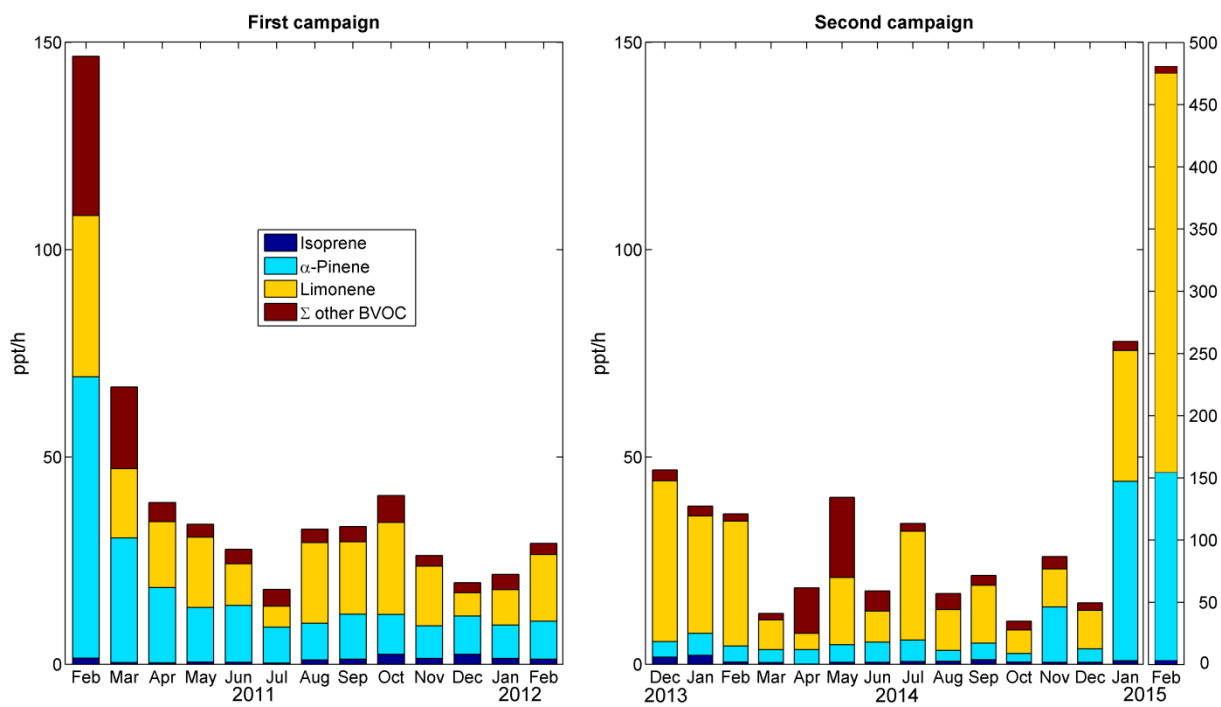
8 Figure 7. Continued.

1 Table 5. Photochemical properties of measured BVOCs during the first and second campaign  
 2 at Welgegund.

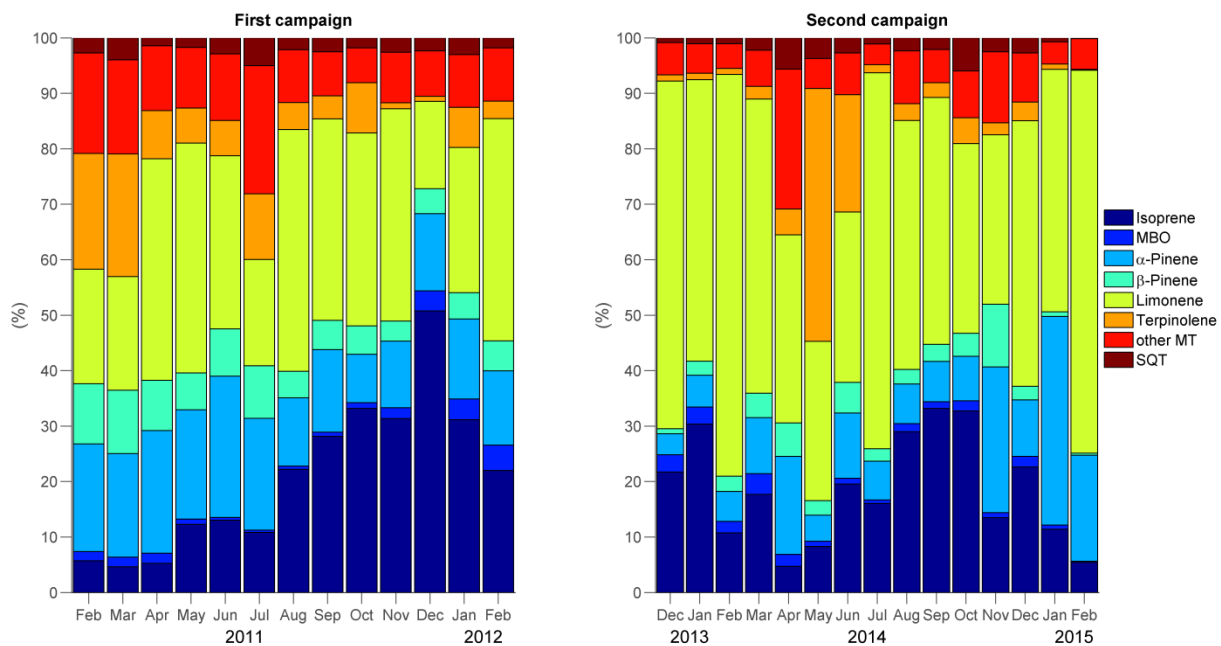
		First period		Second period		[cm <sup>3</sup> molecule <sup>-1</sup> s <sup>-1</sup> ]		
		MIR <sup>a</sup>	Avg	OFP	Avg	OFP	$k_{OH} \times 10^{12}$	$k_{O_3} \times 10^{18}$
	Isoprene	10.28	28	289	23	234	101.0	13.0
	MBO	4.73	12	56	7.7	37	27.5	1.8
	$\alpha$ -Pinene	4.38	71	313	57	251	53.7	86.6
	Camphene		7.9		3.8		53.0	0.9
	$\beta$ -Pinene	3.38	19	64	4.6	16	78.9	15.0
	$\Delta^3$ -Carene	3.13	6.1	19	4.1	13	88.0	37.0
	p-Cymene	4.32	48	206	15	66	15.0	0.05
	1,8-Cineol		13		1.9		22.6	
Monoterpenes	Limonene	4.4	30	131	54	236	171.0	200.0
	Terpinolene	6.16	14	84	28	170	22.5	138.0
	AMCH		6.7		4.2		98.6	430.0
	Nopinene		7.3		11		8.6	
	Bornylacetate		1.7		3.1		7.7	
	4-Allylanisole		11		12		54.3	12.0
	Longicyclene		4.2		1.7		9.4	
	iso-Longifolene		3.0		0.9		96.2	11.4
Sesquiterpenes	Aromadendrene		1.0		2.4		62.5	12.0
	$\alpha$ -Humulene		0.9		2.7		290.0	870.0
	Alloaromadendrene		3.2					

3 <sup>a</sup>MIR denotes maximum incremental reactivity (g O<sub>3</sub>/g VOCs) (Carter, 2009).

4 The rate constants are from Atkinson (2000) and Atkinson and Arey (2003b) except those for  
 5  $\alpha$ -humulene and longifolene OH reaction rates, which are from Shu and Atkinson (1995). Other  
 6 sesquiterpene data is from CSID:1406720, [http://www.chemspider.com/Chemical-  
 7 Structure.1406720.html](http://www.chemspider.com/Chemical-Structure.1406720.html) (last access: 2 May 2016). Predicted data is generated using the US  
 8 Environmental Protection Agency's EPI Suite.



1  
 2 Figure 8a. Monthly means of reaction rates calculated for reactions between O<sub>3</sub> and BVOCs at  
 3 Welgegund. A secondary axis is introduced for reaction rates calculated for February 2015 due  
 4 to much higher reaction rates calculated for this month.



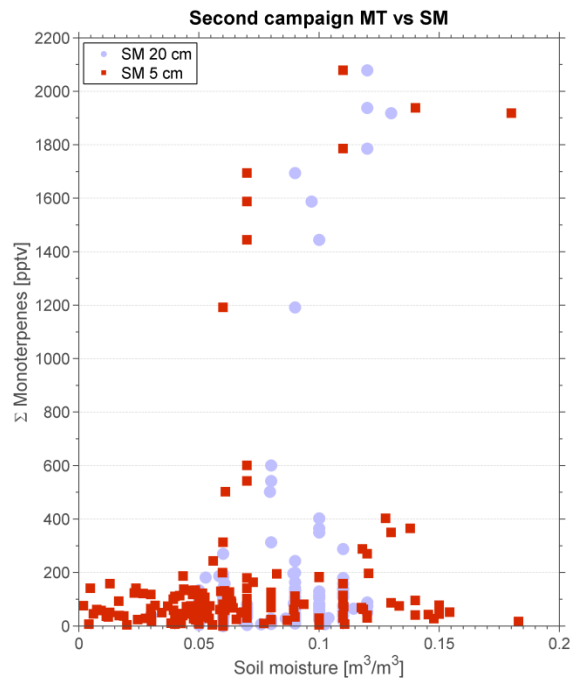
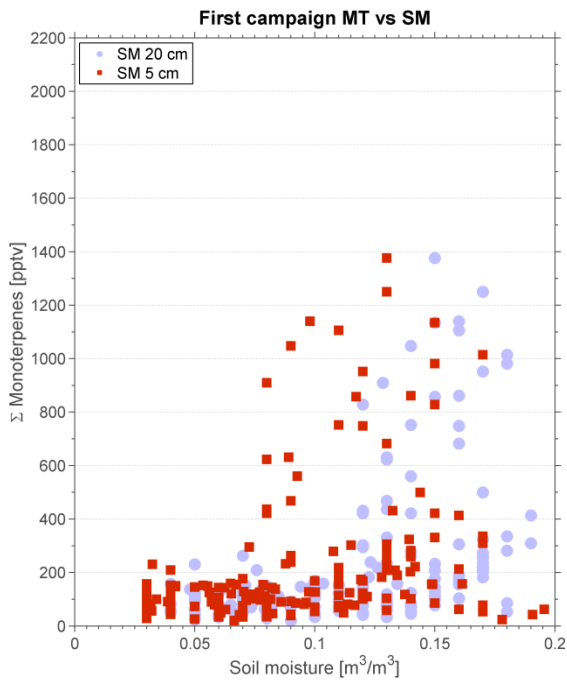
5  
 6 Figure 8b. The relative monthly contribution of different BVOCs to the OH-reactivity at  
 7 Welgegund.  
 8



1 Supplement

2

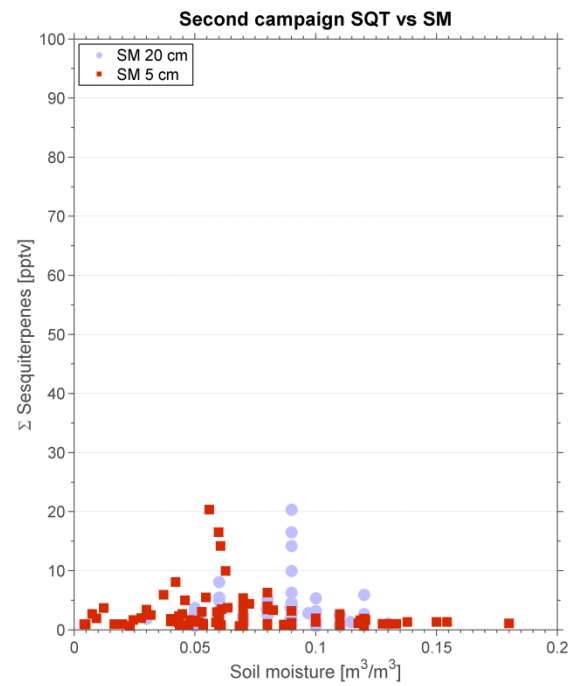
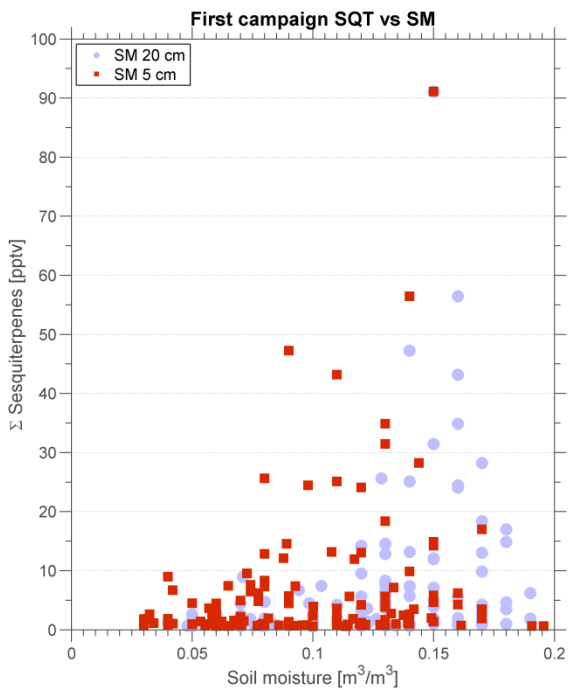
(a)



3

4

(b)



5

6 Figure S1. Correlation between soil moisture and monoterpene concentrations (a), and soil  
7 moisture and SQT (b)