Response to Referee #1

In this manuscript the authors compare results of Master Chemical Mechanism (MCM) simulations of the functional group composition of secondary organic aerosol (SOA) formed from reactions of a-pinene (APIN) and 1,3,5-trimethylbenzene (TMB) with measured values determined by FTIR analysis in order to demonstrate the utility of this approach for evaluating chemical reaction models. The systems chosen for study are popular ones since APIN is an important biogenic VOC and TMB is representative of anthropogenic aromatic emissions. The manuscript is concise and well written, and the authors do a good job of comparing model and measurement results, providing plausible explanations for discrepancies when possible. Because there are in many cases significant differences, and both the simulations and measurements have considerable uncertainties, it is difficult to determine the source of the discrepancies. One conclusion is therefore that more measurements of chamber systems are needed using a variety of tools in order to develop databases of reliable chemical data for model comparisons. In general, however, I think the approach presented here has promise, and that the manuscript presents a useful demonstration of how models can be tested using more detailed chemical data rather than just SOA yields and O/C ratios. I think the paper is suitable for publication in ACP, but have a couple questions the authors should address.

We thank the reviewer for the support and helpful comments. We address specific comments below.

1. **Comment**: The comparison of measurements with simulations of the low-NOx TMB reaction seems problematic, since FTIR does not measure peroxides, which dominate the simulated SOA composition. I don't see any explicit mention of this.

Response: We note that peroxides concentrations were not reported in the study by Sax et al., but is not beyond the capability of analysis by FTIR as there are absorption bands in the infrared window for hydroperoxides and organic peroxides. However, the reviewer's point with respect to model-measurement comparison in this work is worth clarifying in the manuscript. In our pie charts in Figure 3, only the measured mole fraction of OA is shown.

In Methods Section 2.4, we have added the sentence:

"For model-measurement comparison, we select the subset of FGs that are reported by measurement and use relative metrics normalized only by the measured fraction of OA."

2. **Comment**: When certain FG, like peroxides, are not measured by FTIR, how are the reported concentrations of the other FG affected? I do not see an "unidentified" component of SOA in the pie diagrams.

Response In response to the previous comment, we report relative concentrations of the measured fraction. While it is possible to include an unidentified or remaining fraction in the model simulation pie charts, we do not know how large this fraction should be for the FTIR measurements. We have therefore added in the caption of Figure 3 the phrase:

"The mole fractions reported in simulations are summed with respect to the subset of FGs that are reported by measurement to facilitate direct comparison."

3. **Comment**: Is there molecular information available from other experiments conducted under roughly similar conditions that can be used to determine if the major molecular components predicted by the MCM are reasonable, which may help in determining whether the model or measurements are the source of discrepancies in some comparisons?

Response: There are some analyses of gas phase composition which we now include in our analysis. In Results Section 3.1.1:

"Pinonic acid is the second largest contributor to COOH FG, which is consistent with previous reports of pinonic acid being a major contributor to SOA in APIN photooxidation over a range of NO_x conditions Eddingsaas et al. (2012)."

And in Results Section 3.2.1: "As for APIN-lNO $_{\rm x}$, pinonic acid is the second largest contributor to COOH FG; consistent with observations in similar experiments (Eddingsaas et al., 2012)."

We hope that this manuscript will encourage the adoption of FTIR as a complementary tool for OA analysis and that such joint measurements will become available in the future.

4. **Comment**: Are there simpler systems that might be modeled and analyzed?

Response α -pinene dark ozonolysis is a classic, well-studied system (e.g., Yu et al., 1999) which may provide constraints and simplification in interpretation. At present time, we selected simulations for which experimental values were available, but hope that there will be opportunities for further exploration in future studies.

The reasoning for selecting these particular systems have now been explicitly added to the beginning of Methods Section 2:

"We target our model simulations to mimic SOA formation in environmentally controlled chambers for which FG measurements are available."

References

Eddingsaas, N. C., Loza, C. L., Yee, L. D., Seinfeld, J. H., and Wennberg, P. O.: alpha-pinene photooxidation under controlled chemical conditions - Part 1: Gas-phase composition in low- and high-NOx environments, Atmospheric Chemistry and Physics, 12, 6489–6504, doi:10.5194/acp-12-6489-2012, 2012.

Yu, J. Z., Cocker, D. R., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: Gas-phase ozone oxidation of monoterpenes: Gaseous and particulate products, Journal of Atmospheric Chemistry, 34, 207–258, doi:10.1023/A:1006254930583, 1999.

Response to Referee #3

This manuscript introduces a modelling method than can be used for simulating secondary organic aerosol (SOA) formation and for studying especially functional group (FG) contributions in SOA. This kind of modelling approaches are necessary steps towards understanding SOA formation. The modelling framework is reasonable, manuscript is well written and the topic well within the scope of ACP. Therefore I recommend publication of the manuscript in ACP after minor revision.

We thank the reviewer for the positive and helpful comments. We address specific comments below.

- 1. **Comment**: There are number of assumptions/estimations/simplification included in the model and the effect of these on the model results and for model-measurement comparison has not been discussed. Authors state that the relative abundances of functional groups are robust with respect to many of the assumptions (P. 4, L. 23-24). However, I feel this statement needs to be justified. I recommend adding a short discussion/statement on the effect of the following:
 - A. Fixing radiation intensities in the model one constant value (P. 4, L. 6) Does this correspond to conditions in the experiments? If not, are the simulation results sensitive to this?
 - B. A seed is used for initializing the gas-particle partitioning (P. 4, L. 14-15). According to the appendix A the initial seed composition is calculated based on equilibrium gas-particle partitioning theory. Can this initial composition affect the model results at short simulation times? Also, I recommend adding here in the main text a sentence about how the initial seed composition was assigned.
 - C. The vapor wall losses where neglected (P. 4, L. 20). Could this affect the simulations as the wall losses are likely volatility dependent and therefore may affect different FG differently?

Response:

A. The constant radiation intensities are indeed consistent with experimental conditions of Sax et al. (2005) and Chhabra et al. (2011).

In the method section we have therefore added the phrase:

- "Radiation intensities were fixed at their maximum throughout the simulations to mimic conditions used in the chamber studies [...]."
- B. Based on the simulations varying seed composition in Appendix B, the sensitivity to initial seed composition is explored by varying the activity (mole fraction), $a_i = x_i$, according to initial seed concentration. We conclude that after an hour of simulation, relative FG abundances were generally invariant with initial seed composition.

We have added the following statement to Methods Section 2.2:

"We specify the bulk of $C_{\text{OA,init}}$ to be a generic, non-volatile organic solvent that does not participate in reactions or partitioning, and is in equilibrium with the initial composition of the gas-phase (Appendix B). The relative composition reported in this study are insensitive to this value after one hour of simulation (Figures S3–S4)."

We further note in Appendix B that this is in contrast to the approach of (Pankow, 2001):

"Pankow (2001) used $C_{\text{OA,init}}$ and composition (x_i s) from the final C_{OA} values determined from experiments. As this can introduce additional mass into the system, we introduce a different approach."

C. We appreciate the reviewers comment on this point. Under the assumption of a common wall loss parameter (Zhang et al., 2014), it is unlikely that our simulated composition would be altered even while the overall yield simulated would be decreased. However, wall losses that are composition dependent (Matsunaga and Ziemann, 2010; Yeh and Ziemann, 2015) may indeed reduce the most condensible substances in the system and alter the relative particle composition (Cappa et al., 2016; La et al., 2016). However, the magnitude of this effect also depends on the number of condensable species formed and their absolute concentrations. For instance, in the TMB-lNO_x simulation, while gas and particle concentrations are observed to increase, the aerosol composition remains relatively constant. This is possibly due to the fact that the number condensed species is dominated by a few compounds that span a narrow range of saturation concentrations. Also, our model simulations tend to underpredict the oxygenated (polar) fraction in the condensed phase, indicating that the mechanisms for discrepancies suggested in the manuscript are likely to be even more active than current magnitude of differences suggest.

We have included the following comment in Methods Section 2.2:

"However, the impact of vapor losses to chamber walls may require investigation in future work. An assumption of a common wall loss parameter for all species (e.g., Zhang et al., 2014) would mostly reduce the overall yield from simulation, but compound-dependent wall losses (Matsunaga and Ziemann, 2010; Yeh and Ziemann, 2015) may preferentially reduce the concentration of the most condensible substances in the system and lead to a different relative particle composition (Cappa et al., 2016; La et al., 2016). The magnitude of this effect also depends on the number of condensable species formed, the range of saturation concentrations spanned, and their absolute abundance."

2. **Comment**: P. 5, L. 24: 'FTIR analysis has been found to measure around 80 % of the organic mass...' How does the remaining 20 % affect the model-measurement comparison? Could the differences between modelled and measured FG relative contributions be explained in some cases by the missing 20 % of organic mass in measurements?

Response: As the reviewer points out, not all FGs in the simulation that could be present in the physical system are reported by current FTIR calibration models, and this is a likely explanation for underestimation of OA (Russell, 2003). We have decided to remove the statement regarding the quantified fraction as this is more applicable for ambient samples, and is a topic of current study (alluded to in Methods Section 2.4: "Constraining the mapping of measured bonds to atoms for estimation of [mass contributions of functional groups] in various mixtures are planned for future work."). With regards to the model-measurement comparison, only relative fractions considering the measured FGs are shown to circumvent this potential ambiguity.

Also in response to Reviewer #1, We have added a phrase in the Methods Section 2.4 to clarify this point:

"For model-measurement comparison, we select the subset of FGs that are reported by measurement and use relative metrics normalized only by measured fractions of OA."

And in the caption of Figure 3:

"The mole fractions reported in simulations are taken with respect to the sum of FGs that are reported by measurement to facilitate direct comparison."

3. Comment: P. 8, L. 1-2: In addition to missing particle phase aggregation reactions, also highly oxidized high-molecular mass gas phase products might be missing from the model (Ehn et al. 2014, Nature 506, 476-479. Could such highly-oxidized organics explain the discrepancy between the model and measurement?

Response: We thank the reviewer for suggesting this hypothesis. Ehn et al. (2014) report a high mass fraction of (up to two-thirds) by ELVOCs at low to moderate ($\sim 10 \ \mu g \, m^{-3}$) loadings. However, Zhang

et al. (2015) report that ELVOCs may play a smaller role with SVOCs constituting around 60% of the SOA mass at higher loadings, under which the APIN-lNOx and TMB-lNOx simulations fall (Figure S1). However, ELVOCs may play a significant role in the APIN-hNOx and APIN-nNOx scenarios where loadings are smaller. However, the hypothesized mechanism for the formation of ELVOCs are successive H-abstractions and O_2 -additions to peroxy radicals. In the presence of NO_x , reactions of these radicals with NO may prevent significant ELVOC formation.

We have therefore inserted statements in each set of results to discuss the potential role of ELVOCs.

In Results Section 3.1.2:

"This observation suggests that the role played by the gas-phase production of polyfunctional, extremely low volatility compounds (ELVOCs) observed in greater abundance Ehn et al. (2014) at lower aerosol C_{OA} loadings [...]"

In Results Section 3.2.1:

"While production of large, polyfunctional ELVOCs might be a prime candidate for explaining the mass discrepancy at these low $C_{\rm OA}$ loadings (comprising up to two-thirds for mass concentrations less than 10 $\mu \rm g \, m^{-3}$), reactions with NO with peroxy radicals may inhibit formation of ELVOCs through the hypothesized mechanism of H-abstraction and O₂-addition to peroxy radicals (Ehn et al., 2014)."

In Results Section 3.3.2:

"Production and condensation of ELVOCs or additional oxidation mechanisms in the condensed phase not implemented in our model may also contribute to this discrepancy."

In Results Section 3.4.1:

"While formation and condensation of ELVOCs in the experimental system cannot be ruled out, it is likely that their contribution would be much smaller than the SVOC fraction on account of the high mass loadings (Figure S1) (Zhang et al., 2015)."

4. **Comment**: P. 13, L. 19-21: Here it would be good if the authors could summarize also that in which conditions their modelling tool may be applicable and what kind of improvements would be most important ones.

Response: We thank the reviewer for the suggestion to add a general summary.

We have extended the Conclusion Section 4 with the following statements:

"This work illustrates that concurrent measurement of FGs alongside common techniques for atomic and molecular characterization of OA can provide an opportunity for complementary evaluation and further guide detailed understanding of chemical and physical transformations. Analysis of FG abundance can supplement tracking of individual tracers and evaluate the importance of mechanisms that lead to production of a class of compounds in the overall molar (or mass) budget. FG abundances can also provide structural interpretation to variations in elemental ratios (e.g., O:C, H:C, and N:C). Looking forward, systematic model-measurement comparison of FGs under controlled conditions may be able to provide constraints and aid development of chemical mechanism generators (e.g., Gao et al., 2016; Aumont et al., 2005)."

5. **Comment**: Authors mention that for APIN-hNOx and APIN-nNOx cases the model underestimates the SOA mass by an order of magnitude. I was missing this information for the cases APIN-lNOx and TMB-lNOx.

Response: The mass yields were not reported in the APIN-lNO_x and TMB-lNO_x studies as the stated objective of Sax et al. (2005) was to "monitor the temporal change of the functional groups present in the SOA."

Technical comments:

6. **Comment**: P. 7, L. 30-31: 'The sum of the oxidized fraction in the simulation...' This sentence is unclear, please revise.

Response: We have modified this sentence to read: "The oxidized fraction in the simulation is consistently lower than in the measurements, as it remains below 16% in the model while it increases to 40% after 20 hours in the reported measurements."

7. Comment: P. 11, L. 14-26: Why are the model results here compared to Sato et al. measurements although the model simulations were designed to simulate the Sax et al. measurements?

Response: Our simulation is based on Sax et al. (2005) to target comparison of FG measurements. But as Sax et al. (2005) do not report overall O:C ratios in their study, we provide a comparison of simulation results to the overall O:C ratio reported by Sato et al. (2012) for experiments conducted under similar conditions.

We added a modified the phrase to clarify this point:

"The O:C ratio of the simulated SOA is (\sim 0.7). While overall O:C in experiments of Sax et al. (2005) were not reported, Sato et al. (2012) report values in the range of 0.25–0.47 for similar TMB photooxidation experiments with NO_x and methyl nitrate as OH source, with no seed and lower RH (<1%)."

8. **Comment**: Figure 2: Please clarify what 'relative mole fraction compared to the first sample' means. This can be interpreted at least as a ratio between molar fractions of compound i at time t and time t=0s or as moles of compound i at time t divided by moles of compound i at time 0 s.

Response: This is a metric and figure format we adopted from Sax et al. (2005). For clarification, we have added the following statement in the caption of Figure 2: "For a chosen FG the changes of the relative mole fraction compared to the first sample is calculated as the ratio between the relative mole fraction at the chosen time and the relative mole fraction at 1 hour."

9. Comment: Figure 2: Please clarify if the naCO includes both ketone and aldehyde for the model or if naCO includes only aldehyde for the model as ketone is presented also separately. At the moment it's not clear in measured naCO should be compared to modelled naCO or the sum of modelled naCO and ketone.

Response: We thank the reviewer for this opportunity for clarification. In the Figure 2 caption, we now write: "naCO includes ketone and aldehyde FGs, but the change in relative ketone FG abundance is also shown separately for illustration."

10. **Comment**: Figure 2: Consider plotting a vertical line at [y] = 1.0. That would make it easier to read the figure in case of those compounds that show only small changes.

Response: The dashed line has been added to the figure and we have added also a phrase in the caption: "The dashed line corresponds to y=1 and has been added for visual reference."

References

Aumont, B., Szopa, S., and Madronich, S.: Modelling the evolution of organic carbon during its gas-phase tropospheric oxidation: development of an explicit model based on a self generating approach, Atmospheric Chemistry and Physics, 5, 2497–2517, doi:10.5194/acp-5-2497-2005, 2005.

Cappa, C. D., Jathar, S. H., Kleeman, M. J., Docherty, K. S., Jimenez, J. L., Seinfeld, J. H., and Wexler,
 A. S.: Simulating secondary organic aerosol in a regional air quality model using the statistical oxidation
 model – Part 2: Assessing the influence of vapor wall losses, Atmospheric Chemistry and Physics, 16,
 3041–3059, doi:10.5194/acp-16-3041-2016, 2016.

- Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan, R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic aerosol, Atmospheric Chemistry and Physics, 11, 8827–8845, doi:10.5194/acp-11-8827-2011, 2011.
- Ehn, M., Thornton, J. A., Kleist, E., Sipil'a, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Pet'aj'a, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479, doi:10.1038/nature13032, 2014.
- Gao, C. W., Allen, J. W., Green, W. H., and West, R. H.: Reaction Mechanism Generator: Automatic construction of chemical kinetic mechanisms, Computer Physics Communications, 203, 212 225, doi: http://dx.doi.org/10.1016/j.cpc.2016.02.013, 2016.
- La, Y. S., Camredon, M., Ziemann, P. J., Valorso, R., Matsunaga, A., Lannuque, V., Lee-Taylor, J., Hodzic, A., Madronich, S., and Aumont, B.: Impact of chamber wall loss of gaseous organic compounds on secondary organic aerosol formation: explicit modeling of SOA formation from alkane and alkene oxidation, Atmospheric Chemistry and Physics, 16, 1417–1431, doi:10.5194/acp-16-1417-2016, 2016.
- Matsunaga, A. and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield Measurements, Aerosol Science and Technology, 44, 881–892, doi:10.1080/02786826.2010.501044, 2010.
- Pankow, J. F.: A consideration of the role of gas/particle partitioning in the deposition of nicotine and other tobacco smoke compounds in the respiratory tract, Chemical Research In Toxicology, 14, 1465–1481, doi:10.1021/tx0100901, 2001.
- Russell, L. M.: Aerosol organic-mass-to-organic-carbon ratio measurements, Environmental Science & Technology, 37, 2982–2987, doi:10.1021/es026123w, 2003.
- Sato, K., Takami, A., Kato, Y., Seta, T., Fujitani, Y., Hikida, T., Shimono, A., and Imamura, T.: AMS and LC/MS analyses of SOA from the photooxidation of benzene and 1,3,5-trimethylbenzene in the presence of NO_x: effects of chemical structure on SOA aging, Atmospheric Chemistry and Physics, 12, 4667–4682, doi:10.5194/acp-12-4667-2012, 2012.
- Sax, M., Zenobi, R., Baltensperger, U., and Kalberer, M.: Time resolved infrared spectroscopic analysis of aerosol formed by photo-oxidation of 1,3,5-trimethylbenzene and alpha-pinene, Aerosol Science and Technology, 39, 822–830, doi:10.1080/02786820500257859, 2005.
- Yeh, G. K. and Ziemann, P. J.: Gas-Wall Partitioning of Oxygenated Organic Compounds: Measurements, Structure–Activity Relationships, and Correlation with Gas Chromatographic Retention Factor, Aerosol Science and Technology, 49, 727–738, doi:10.1080/02786826.2015.1068427, 2015.
- Zhang, X., Schwantes, R. H., Coggon, M. M., Loza, C. L., Schilling, K. A., Flagan, R. C., and Seinfeld, J. H.: Role of ozone in SOA formation from alkane photooxidation, Atmospheric Chemistry and Physics, 14, 1733–1753, doi:10.5194/acp-14-1733-2014, 2014.
- Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.: Formation and evolution of molecular products in α -pinene secondary organic aerosol, Proceedings of the National Academy of Sciences, 112, 14168–14173, doi:10.1073/pnas.1517742112, 2015.

Model-measurement comparison of functional group abundance in α -pinene and 1,3,5-trimethylbenzene secondary organic aerosol formation

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Abstract. Secondary organic aerosol (SOA) formed by α -pinene and 1,3,5-trimethylbenzene photooxidation under different NO_x regimes is simulated using the Master Chemical Mechanism v3.2 (MCM) coupled with an absorptive gas/particle partitioning module. Vapor pressures for individual compounds are estimated with the SIMPOL.1 group contribution model for determining apportionment of reaction products to each phase. We apply chemoinformatic tools to harvest functional group (FG) composition from the simulations and estimate their contributions to the overall oxygen to carbon ratio. Furthermore, we compare FG abundances in simulated SOA to measurements of FGs reported in previous chamber studies using Fourier Transform Infrared Spectroscopy. These simulations qualitatively capture the dynamics of FG composition of SOA formed from both α -pinene and 1,3,5-trimethylbenzene in low NO_x conditions, especially in the first hours after start of photooxidation. Higher discrepancies are found after several hours of simulation; the nature of these discrepancies indicate sources of uncertainty or types of reactions in the condensed or gas phase missing from current model implementation. Higher discrepancies are found in the case of α -pinene photooxidation under different NO_x concentration regimes, which are reasoned through the domination by a few polyfunctional compounds that disproportionately impact the simulated FG abundance in the aerosol phase. This manuscript illustrates the usefulness of FG analysis to complement existing methods for model-measurement evaluation.

1 Introduction

Atmospheric aerosols are complex mixtures that can contain a multitude of chemical species (Seinfeld and Pandis, 2006). While the inorganic fraction comprises a relatively small number of compounds, the organic fraction (or organic aerosol, OA) includes thousands of compounds with diverse molecular structures (Hamilton et al., 2004). These compounds take part in multitude of gas phase, aerosol phase, and heterogeneous transformation processes (e.g., Kroll and Seinfeld, 2008; Hallquist et al., 2009; Ziemann and Atkinson, 2012) that must be modeled with sufficient fidelity to predict atmospheric concentrations and impacts from various emission scenarios.

A mechanism central to these processes is the formation of semivolatile organic compounds (SVOCs) through gas-phase oxidation of volatile organic compounds (VOC) precursors and their reaction products. α -pinene (APIN) and 1,3,5-trimethylbenzene (TMB) are examples of biogenic and anthropogenic VOC precursors, respectively, which have been studied for their chemical

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reaction mechanisms and aerosol yields in environmentally-controlled chamber experiments and numerical simulation. APIN is a monoterpene compound primarily emitted from coniferous vegetation (Fuentes et al., 2000; Tanaka et al., 2012) with high emission rate, reactivity, and secondary organic aerosol (SOA) generation potential (e.g., Fehsenfeld et al., 1992; Lamb et al., 1993; Chameides et al., 1988; Jenkin, 2004; Tolocka et al., 2004; Sindelarova et al., 2014). TMB is an aromatic compound emitted from vehicular emissions and a major contributor to urban organic aerosol (e.g., Kalberer et al., 2004); its degradation mechanism has also been subject of collective evaluation (Metzger et al., 2008; Wyche et al., 2009; Rickard et al., 2010; Im et al., 2014). Gas-phase oxidation reactions are modeled with chemically explicit or semi-explicit treatment, or alternatively using a basis set approach based on simplified molecular or property descriptors; SOA formation is commonly modeled by coupling these reactions with partitioning of oxidation products to an absorptive organic phase (e.g., Jenkin et al., 1997; Pun et al., 2002; Griffin et al., 2003; Aumont et al., 2005; Capouet et al., 2008 Figgans et al., 2010; Barley et al., 2011; Chen et al., 2011; Murphy et al., 2011; Aumont et al., 2012; Jathar et al., 2015; McVay et al., 2016). SVOCs produced by such reactions can in reality partition among multiple phases (vapor, organic liquid, aqueous, solid), and participate in additional functionalization, accretion, or fragmentation reactions in one of many phases (Kroll et al., 2011; Cappa and Wilson, 2012; Im et al., 2014; Zhang and Seinfeld, 2013; Zhang et al., 2015). These processes are represented in models with varying degrees of detail; simplifying or wholly omitting various mechanisms out of concerns for computational feasibility or lack of sufficient knowledge. For instance, in a work we follow closely in this manuscript, Chen et al. (2011) used a fully explicit gas-phase reaction mechanism with absorptive organic partitioning and evaluated the potential importance of missing heterogeneous and condensed-phase mechanisms based on discrepancy of model simulation and experiments.

Our capability to simulate SOA formation is often evaluated against aerosol mass yield, O:C, carbon oxidation state, mean carbon number, volatility, and specific species or compound classes when available (e.g., Robinson et al., 2007; Kroll et al., 2011; Donahue et al., 2012; Nozière et al., 2015). These properties can be measured using various forms of mass spectrometry (e.g., Jayne et al., 2000; Jimenez et al., 2009; Nizkorodov et al., 2011); or monitoring changes in size distribution in combination with isothermal dilution or thermal heating (e.g., Grieshop et al., 2009; Cappa, 2010; Epstein and Donahue, 2010; Donahue et al., 2012). Functional group (FG) composition is a complementary representation of organic molecules and complex organic mixtures that offers a balance between parsimony and chemical fidelity for measurement and interpretation.

FGs represent structural units of molecules that play a central role in chemical transformations, and provide insight into evolution of complex organic mixtures without monitoring all species explicitly (Holes et al., 1997; Sax et al., 2005; Presto et al., 2005; Lee and Chan, 2007; Chhabra et al., 2011; Zeng et al., 2013). FG abundances have also been associated with volatility (e.g., Pankow and Asher, 2008), hygroscopicity (e.g., Hemming and Seinfeld, 2001; Suda et al., 2014), and magnitude of nonideal interactions in the condensed phase (e.g., Ming and Russell, 2002; Griffin et al., 2003; Zuend et al., 2011). However, two impediments have been the likely cause of slow adoption of this representation. Building quantitative calibration models of FG abundance have posed analytical challenges, but rapid progress has been made over the past decade with Fourier Transform Infrared Spectroscopy (FTIR) (e.g., Sax et al., 2005; Reff et al., 2007; Coury and Dillner, 2008; Day et al., 2010; Takahama et al., 2013; Ruthenburg et al., 2014; Takahama and Dillner, 2015), Nuclear Magnetic Resonance (Decesari et al., 2007; Cleveland et al., 2012), spectrophotometry (Aimanant and Ziemann, 2013; Ranney and Ziemann, 2016), and GC-MS with

derivatization (Dron et al., 2010). The second challenge is computationally harvesting FG abundance from a large set of known molecular structures. To this end, Ruggeri and Takahama (2015) developed a set of substructure definitions corresponding to FGs that can be queried against arbitrary molecules specified by their molecular graphs.

In this work, we apply these new substructure definitions to describe the FG composition of products simulated by gas-phase reactions prescribed with the Master Chemical Mechanism (MCMv3.2) (Jenkin et al., 1997; Saunders et al., 2003; Jenkin et al., 2003; Bloss et al., 2005), and SOA constituents formed by their dynamic absorptive partitioning (Chen et al., 2011). Three instances of APIN photooxidation under varying initial concentrations of oxides of nitrogen (NO_x), and TMB oxidation in the presence of NO_x are studied in accordance to aerosol FG composition characterized by Sax et al. (2005) and Chhabra et al. (2011) in chamber studies using FTIR. The model results are analyzed through a suite of FG abundances and model-measurement comparisons of measured FGs are presented to hypothesize reasons (including unimplemented mechanisms) for discrepancies where they occur.

2 Methods

We target our model simulations to mimic SOA formation in environmentally controlled chambers for which FG measurements are available.

15 2.1 Systems studied

Photooxidation of APIN under "low NO_x " (NO_x /APIN of 0.8), "high NO_x " (NO_x /APIN of 18), and no NO_x conditions (designated as INO_x), h NO_x , and n NO_x , respectively), and TMB under "low NO_x " (NO_x /TMB ratio of 0.24; designated as INO_x) conditions were simulated in this study to compare with available measurements of aerosol FG composition in environmental chamber experiments. Simulations were run at 298 K and conditions closely following experimental descriptions summarized in Table 1, with a few exceptions. In the case of APIN degradation in high NO_x conditions, the H_2O_2 was used as the OH radical initiator as CH_3ONO is not available in the MCMv3.2 degradation scheme. When the reacted instead of initial precursor concentration is reported, this value is used as the initial concentration for the simulations. This decision is supported by the virtual observation that 99% of the precursor is reacted after 4.5–6.5 hours in these cases (Figure S1), and specification of higher initial concentrations lead to reacted quantities inconsistent with experimental specifications.

25 2.2 Model formulation

While differing in implementation, the model specification resembles the MCM-SIMPOL framework described by Chen et al. (2011). The chemical mechanism prescribed by MCMv3.2 (Jenkin et al., 1997; Saunders et al., 2003; Jenkin et al., 2003; Bloss et al., 2005) was used to simulate the gas-phase oxidation of volatile organic compounds (VOCs). The Kinetic Pre-Processor (KPP, Damian et al., 2002; Sandu and Sander, 2006; Henderson, 2016) was used to generate the gas-phase chemistry code in Fortran 90. A separate dynamic absorptive partitioning (Pankow, 1994) module was added via sequential operator splitting (Yanenko, 1971; Orlan and Boris, 2000; Vayenas et al., 2005) to simulate gas/particle (G/P) partitioning after the reaction

operator. Pure component vapor pressures of organic compounds in the MCMv3.2 degradation schemes were calculated using SIMPOL.1 (Pankow and Asher, 2008), and non-ideal interactions were neglected in these simulations (i.e., activity coefficients were set to unity for all species). Vapor pressures are converted to equivalent mass concentrations C^0 (Appendix A), and normalized by a reference value for presentation in logarithmic units (Seinfeld and Pandis, 2006) such that the notation $\log C^0$ implies $\log_{10}(C^0/1\,\mu\mathrm{g\,m^{-3}})$. LSODE (Livermore Solver for Ordinary Differential Equations; Radhakrishnan and Hindmarsh, 1993) was used as the numerical solver for each operation (reaction and G/P partitioning). A time step of 60 seconds is used in this study, as it is in the order of magnitude of the timescale of gas-phase oxidation and condensation/evaporation under chamber conditions (Cocker et al., 2001) and leads to stable solutions. Radiation intensities were fixed at their maximum throughout the simulations to mimic conditions used in the chamber studies, with values corresponding to clear sky conditions at an altitude of 0.5 km, 1° solar zenith angle in July, and a latitude of 45 N (Derwent et al., 1996; Hayman, 1997; Derwent et al., 1998; Saunders et al., 2003).

Absorptive partitioning to a purely organic phase is considered in this model (Appendix A in Supplemental Material). The relative humidity (RH) specified in the experiments are converted to equivalent concentrations of H_2O for participation in the HO₂ radical self reaction to form hydrogen peroxide (Mozurkewich and Benson, 1985), but water uptake by the aerosol and its influence on G/P partitioning of organic compounds (Seinfeld et al., 2001; Chang and Pankow, 2010) is not considered. As aerosol growth following homogeneous and heterogeneous nucleation processes of the condensed organic phase in the chamber experiments are not included in the model, we use a seed $C_{OA,init}$ of 1 $\mu g \, m^{-3}$ to initiate G/P partitioning (Appendix B). We specify the bulk of $C_{OA,init}$ to be a generic, non-volatile organic solvent that does not participate in reactions or partitioning, and is in equilibrium with the initial composition of the gas-phase (Appendix B). The relative composition reported in this study are insensitive to this value after one hour of simulation (Figures S3-S4). To differentiate between the SOA formed in the simulation and the total organic aerosol phase involved in partitioning, we denote the former quantity as C_{SOA} the latter as $C_{\rm OA} = C_{\rm OA,init} + C_{\rm SOA}$. No condensed-phase reactions are included; as with Chen et al. (2011) we consider them as a potential source of model-measurement discrepancies. While the particle diameter of the monodisperse population is allowed to grow according to the organic aerosol condensed (Section B), the number concentration of particles is kept fixed during the simulation; losses of both particles and gases to chamber walls (e.g., Loza et al., 2010; Matsunaga and Ziemann, 2010; Zhang et al., 2014a) are neglected. These assumptions will affect calculations of total yield and rate of change in aerosol mass; however, aerosol mass yields are in the range of physical expectation (Figure S5; mass concentrations represented in the volatility basis set convention are also shown in Figure S6 for reference). Relative abundances of functional groups are robust with respect to many of these assumptions and will be the primary focus of our presentation and model-measurement comparisons. However, the impact of vapor losses to chamber walls may require investigation in future work. An assumption of a common wall loss parameter for all species (e.g., Zhang et al., 2014b) would mostly reduce the overall yield from simulation, but compound-dependent wall losses (Matsunaga and Ziemann, 2010; Yeh and Ziemann, 2015) may preferentially reduce the concentration of the most condensible substances in the system and lead to a different relative particle composition (Cappa et al., 2016; La et al., 2016). The magnitude of this effect also depends on the number of condensable species formed, the range of saturation concentrations spanned, and their absolute abundance.

2.3 Simulation analysis

A chemoinformatic tool (APRL-SSP; Takahama, 2015) described by Ruggeri and Takahama (2015) is used to harvest FG abundances (enumeration of the FG fragments) from each molecule in the simulations. This tool consists of scripts invoking OpenBabel and Pybel (O'Boyle et al., 2008, 2011) and SMARTS patterns (DAYLIGHT Chemical Information Systems, Accessed 30 September 2015) formulated and validated for these chemical systems. Using this tool, molecular structure is mapped to input parameters for SIMPOL.1, and FG abundances of the organic aerosol mixture are obtained from molecular concentrations. Most importantly, we extract two arrays with elements ϕ_{ip} , the number of times FG p occurs in molecule i, and ϕ_{ipa}^* , the number of times atom type a occurs in FG p in molecule i. We combine these two coefficients with the molecular or molar concentrations C of compound i in phase α generated by our simulations to estimate several useful mixture properties for time t_j :

$$\begin{split} \sum_{i \in \mathcal{M}} C_i^\alpha(t_j) \, \phi_{ip} &= \text{abundance of FG } p \\ C_i^\alpha(t_j) \, \phi_{ip} / \left(\sum_{i \in \mathcal{M}} C_i^\alpha(t_j) \, \phi_{ip} \right) &= \text{fractional contribution of molecule } i \text{ to abundance of FG } p \\ \sum_{i \in \mathcal{M}} C_i^\alpha(t_j) \, \phi_{ipa}^* &= \text{apportionment of atoms of type } a \text{ to FG } p \;. \end{split}$$

The summation is taken for the set of all compounds (or molecule types) \mathcal{M} . The last quantity is used to separate the contributions of O:C and N:C from various FGs. The set of patterns were constructed to meet conditions of completeness and specificity (each atom is matched only by one and only one group) such that the sum of oxygen and nitrogen atoms in each FG sums to the total number of atoms in the system (Ruggeri and Takahama, 2015). Polyfunctional carbon atoms are not considered in the condition for specificity — matches by multiple groups leads to overestimation of counts in ϕ_{ipa}^* — therefore, the total number of carbon used in the denominator of these atomic ratios is estimated using the SMARTS pattern [#6].

We additionally estimate Integrated Reaction Rates (IRR; Jeffries and Tonnesen, 1994) to examine degradation rates relative to rates of production in the gas phase (g) for selected systems. IRR for reaction r affecting compound i at time t_j is calculated from the rate constant k and the product of concentrations C:

$$IRR_{ri}(t_j) = C_i^{(g)}(t_j) - C_i^{(g)}(t_j - \Delta t) = \int_{t_j - \Delta t}^{t_j} dt \left(k_r \prod_{i' \in \mathcal{M}_r} C_{i'}^{(g)}(t) \right)$$
$$\approx \Delta t \left(k_r \prod_{i' \in \mathcal{M}_r} C_{i'}^{(g)}(t_j) \right).$$

5 \mathcal{M}_r is the set of compounds involved in reaction r. The expression in parentheses is the conventional rate equation for reaction r. To obtain the IRR for functional group p, we multiply by the factor ϕ_{ip} described above:

$$IRR_{rp}(t_j) = \sum_{i \in \mathcal{M}_r} IRR_{ri}(t_j) \phi_{ip} .$$

IRR estimates were harvested from the LSODE solver, and the PERMM package (Henderson, 2015) used to associate compounds and FGs with each reaction.

2.4 Measurements

FTIR analysis reported by Sax et al. (2005) and Chhabra et al. (2011) quantified the molar abundance of alkane CH (aCH), carboxylic acid (COOH), non-acid (ketone and aldehyde) carbonyl (naCO), alcohol OH (aCOH), and organonitrate (CONO₂) FGs. Uncertainties in the FG quantification have been reported to be between 5 and 30% (Russell, 2003; Takahama et al., 2013). Sax et al. (2005) collect particles in the range of 86-343 nm onto zinc selenide substrates by impaction, while Chhabra et al. (2011) sample generated aerosol onto Polytetrafluoroethylene (PTFE) filters for FTIR analysis. Measurement artifacts can arise during time-integrated collection of aerosol samples and can differ according to duration of sampling (Subramanian et al., 2004) or method of collection (Zhang and McMurry, 1987). The primary driver for absorptive and evaporative artifacts which may impact bulk mass estimation is the difference between the changing gas-phase composition and equilibrium vapor composition with respect to the aerosol phase, but model simulations suggest the relative gas-phase composition stabilizes after the first few hours. Changes in particle composition due to condensed-phase chemistry may perturb the equilibrium, but this phenomenon may be interpreted together with condensed-phase processes not included in the model. In the analysis of Chhabra et al. (2011), samples transported off-site for analysis were frozen to minimize evaporation and reaction artifacts during storage. Additionally, evaporative losses in the analysis chamber of the FTIR (during purging of headspace with dry nitrogen gas) were minimized by rapid scanning, and Sax et al. (2005) report that the spectrum was stable even when repetitive measurements are performed.

In this work, we limit our discussion to results based on molar rather than mass concentrations of FG abundances. While mass concentrations are commonly reported for FTIR measurements of ambient samples (e.g., Russell et al., 2009), estimates are based on fixed assumptions regarding the apportionment of polyfunctional carbon atoms to associated FGs (e.g., Allen et al., 1994; Russell, 2003; Reff et al., 2007; Takahama et al., 2013; Ruthenburg et al., 2014). These assumptions can affect both mass estimation and atomic ratios (e.g., O:C). Chhabra et al. (2011) proposed a modification based on assumed molecular structures in their chamber experiments, and mass estimates using these values are shown in Figure S7. Constraining the mapping of measured bonds to atoms for estimation of these quantities in various mixtures are planned for future work. For model-measurement comparison, we select the subset of FGs that are reported by measurement and use relative metrics normalized only by measured fractions of OA.

3 Results and Discussion

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In each of the following sections, we begin by describing the simulated evolution of FGs primarily in terms of their contribution to the O:C ratio (Figure 1), and then discuss comparisons of mole fractions with observations for a subset of measured FGs (Figures 2 and 3).

3.1 APIN-INO_v

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3.1.1 Simulation results

Initially, only the most oxygenated species condense to the aerosol phase, but oxygenated products continue to be formed in the gas phase and the O:C values exceeds the aerosol-phase O:C after four hours. The O:C ratios approach 0.75 and 0.6 for the gas and aerosol phases, respectively, after 20 hours of simulation (Figure 1). The O:C ratio of the simulated aerosol phase is comparable to the O:C ratio measured by Chen et al. (2011) and Zhang et al. (2015) in ozonolysis and photooxidation experiments without NO_x (\sim 0.5 in both cases).

The FG that contributes the most to the aerosol O:C ratio after 20 hours is hydroperoxide (31%), while in the gas phase the peroxyacyl nitrate is the major contributor (carrying five oxygen atoms per peroxyacyl nitrate FG) with 55% of the O:C ratio of the gas phase mixture. Some peroxyacyl nitrates are also partitioned to the aerosol phase as reported in laboratory measurements (Jang and Kamens, 2001), but make a smaller contribution (12%) to the aerosol O:C. aCOH and CONO₂ FGs are found in higher abundance in the aerosol phase than many other FGs (Section S4) and contribute to the aerosol-phase O:C, while contributing negligibly to the gas-phase O:C. The large contribution of hydroperoxide FG to the aerosol-phase O:C is consistent with their large contributions to SOA mass suggested in previous studies (Bonn et al., 2004; Wang et al., 2011; Mertes et al., 2012).

Addition of COOH lowers the pure component vapor pressure of a given molecule by four orders of magnitude (Kroll and Seinfeld, 2008; Pankow and Asher, 2008), but contribution to gas and aerosol phase O:C are approximately equal. In the gasphase, CH3CO2H (formed from degradation of the peroxyacid radical compound CH3CO3) constitutes 60% of the COOH fraction (Figure 4) at maximum C_{SOA} (9.3 hours). The aldehyde and ketone CO lower the pure component vapor pressure of around one order of magnitude (Kroll and Seinfeld, 2008), but their contribution to O:C is greater than COOH in the aerosol phase on account of the higher abundance of carbonyl-containing compounds. More than 80% of the moles of carbonyl in both the gas and aerosol phases are associated with ketone rather than aldehyde CO (Figure S8).

We note the prevalence of several large polyfunctional compounds contributing to aerosol-phase. Their cumulative contributions to the total abundance varies over time (Figures S9 and S10); their contributions at peak $C_{\rm SOA}$ are shown in Figure 4. Four compounds (C97OOH, C98OOH, C106OOH, and C719OOH) comprise 70% of the ketone and 80% of the hydroperoxide abundance. 811PAN contributes 50% of the peroxyacyl nitrate and also 45% of the COOH. Illustrations for these compounds are provided in Table 2. Pinonic acid is the second largest contributor to COOH FG, which is consistent with previous reports of pinonic acid being a major contributor to SOA in APIN photooxidation over a range of NO_x conditions Eddingsaas et al. (2012).

3.1.2 Model-measurement comparison of FG mole fractions

Qualitative changes in the mole fractions of COOH, aCOH and CONO₂ FGs over the initial values reported by Sax et al. (2005) is well captured by the model (Figure 2). The magnitude of increase in COOH is higher in the measurements than in the model: an increase of 2.6 times against 1.5 times can be seen between the beginning and the end of the measurements and

the simulation, respectively. For aCOH the discrepancy is smaller; an increase of 1.5 times from the beginning to the end of the experiments against 1.3 in the simulation is found. For CONO₂, the relative mole fraction decreases from 1 to 0.6 during the experiment, while the model predicts a decrease to 0.3. For carbonyl (CO), the model is able to capture the general trend of initial decrease followed by an increase after 4 hours. The trend in modeled naCO is largely contributed by ketone, as it comprises more than 80% of the naCO (Figure S8). The magnitude of decrease in relative mole fraction of aCH observed by Sax et al. (2005) is not captured by the model. The measured relative mole fraction compared to the first sample decreases from 1 to 0.8, while its change is not detectable in the simulation (Figure 2).

The evolving differences in mole fractions between measurement and model are better viewed in Figure 3. The oxidized fraction in the simulation is consistently lower than in the measurements, as it remains below 16% in the model while it increases to 40% after 20 hours in the reported measurements. We consider two condensed-phase reaction mechanisms that may lead to such differences. Viewing the distribution of the compounds present in the MCM APIN-INO_x degradation scheme on $\log C^0$ vs. molar mass space (Figure 5), we see that the model does not include lower volatility compounds with molecular masses higher than 300 g mol⁻¹ observed in experiments (Shiraiwa et al., 2014). This high molecular mass fraction cannot entirely explain the missing COOH, aCOH, and naCO, however, as accretion reactions do not significantly increase the O:C of the mixture (Shiraiwa et al., 2014; Zhang et al., 2015). In the analysis by Shiraiwa et al. (2014), these compounds with high molecular mass and low volatility have an O:C ratio between 0.3 and 0.6. Furthermore, Zhang et al. (2015) report that around 60% of the APIN SOA mass generated in environmental controlled chamber experiments for loadings () is constituted by SVOCs. This observation suggests that the role played by the gas-phase production of polyfunctional, extremely low volatility compounds (ELVOCs) observed in greater abundance Ehn et al. (2014) at lower aerosol C_{OA} loadings and condensed-phase dimerization reactions can only partly be responsible for the discrepancies between simulations and experiments that we report in this study. Proposed dimerization reactions do not contribute to depletion of aCH bonds, and dimers produced in the aerosol phase have been found to have similar O:C ratio to the monomer (Zhang et al., 2015). Photolysis of hydroperoxides has been suggested as a condensed phase mechanism that leads to increase in naCO (Epstein et al., 2014), but an estimate based on the 6-day lifetime molar conversion of hydroperoxide groups to naCO only increases the latter fraction from 8% to 9% of the FG mole fraction after 21 hours (though naCO increases by 9% over the case of no conversion), and does not fully explain the discrepancy between model and measurements for this FG. Further oxidation due to dissolved oxidants, such as OH radical, however may reduce the proportion of aCH relative to oxidized groups, though this rate is also dependent on diffusion and uptake of these radicals by the SOA (Donahue et al., 2013).

3.2 APIN-hNO_x

3.2.1 Simulation results

While the FGs present in APIN-hNO_x system are identical to the APIN-lNO_x system, we find they occur in different proportions on account of both the lower ratio of VOC precursor to NO_x concentrations, and lower absolute precursor concentrations. The predicted aerosol O:C ratio in this simulated system is approximately 0.75, while Chhabra et al. (2011) reports experimental

values around 0.4 according to AMS measurements. $CONO_2$ accounts for 47% of the simulated aerosol O:C after 20 hours (Figure 1). Both aldehyde and ketone CO contribute to O:C in the gas-phase more than in the aerosol phase, while $CONO_2$, aCOH, and COOH contribute primarily to O:C in the aerosol phase. The predicted aerosol N:C ratio is also overestimated (\sim 0.1 in the simulated aerosol, Figure S11) compared to the measured value of 0.03, on account of the large contribution from $CONO_2$.

Lower precursor concentrations in the Caltech chamber experiments (Table 1) lead to lower concentration of condensible products in these corresponding simulations (Figure S1), enabling only a few compounds to partition to the aerosol phase in significant quantities (Figures S6, S9, and S10). The aerosol fraction of COOH exceeds 10%, but the rest remain below 5% of the gas phase, in contrast to the APIN-INO_x system where the aerosol fraction of six FGs exceed 10% (Figure S12). The aerosol mass yields on the order of a few percent (Figure S5) are consistent with C_{SOA} produced in the presence of high NO_x concentrations (e.g., Ng et al., 2007), where NO can compete for reaction with peroxy radicals that may otherwise produce lower volatility products. However, overall C_{SOA} formed is an order of magnitude lower than the 54 $\mu g m^{-3}$ reported in the experiments (Chhabra et al., 2011), which is surprising given that chamber experiments without wall loss corrections tend to underestimate true yields (e.g., Zhang et al., 2014a). This underprediction may suggest the increasing importance of oligomer formation (e.g., Gao et al., 2004; Tolocka et al., 2004; Kalberer et al., 2006; Kroll and Seinfeld, 2008; Chen et al., 2011; Chhabra et al., 2011; Hall and Johnston, 2011) relative to the absorptive partitioning pathway at low C_{OA} concentrations (Presto and Donahue, 2006). While production of large, polyfunctional ELVOCs might be a prime candidate for explaining the mass discrepancy at these low $C_{\rm OA}$ loadings (comprising up to two-thirds for mass concentrations less than $10~{\rm \mu g}~{\rm m}^{-3}$), reactions with NO with peroxy radicals may inhibit formation of ELVOCs through the hypothesized mechanism of H-abstraction and O₂-addition to peroxy radicals (Ehn et al., 2014). Sensitivity analyses conducted to increase the rate of condensation and overall C_{SOA} formed had little impact on relative abundances estimated for FGs (Appendix B), so the interpretations presented are robust for the gas-phase reaction mechanisms included and vapor pressures prescribed in our simulations.

In Figure 6, we see that C813NO3 is a polyfunctional compound that comprises 75% of CONO₂, 95% of COOH, and 70% ketone CO abundance, and 75% of the $C_{\rm SOA}$ mass in the simulated aerosol at peak $C_{\rm SOA}$ (3.2 hours). As for APIN-INO_x, pinonic acid is the second largest contributor to COOH FG; consistent with observations in similar experiments (Eddingsaas et al., 2012). Polyfunctionality may introduce challenges in vapor pressure for linear group contribution methods such as SIM-POL.1, so we evaluate the uncertainty in vapor pressure prediction of the top five contributors (C813NO3, C98NO3, C719NO3, APINANO3, and APINBNO3) to the CONO₂ abundance and $C_{\rm SOA}$ mass by comparing to other methods (Table 3). SIMPOL.1 has been found to generally predict lower vapor pressures compared to other estimation methods like EVAPORATION (Compernolle et al., 2011) and the method of Nannoolal (Nannoolal et al., 2008), but in the case of mononitrates Compernolle et al. (2011) report that differences with EVAPORATION and the Myrdal-Yalkowsky method (Myrdal and Yalkowsky, 1997) are negligible. For these critical compounds, the vapor pressures estimated by SIMPOL.1 are in the range of other estimates except for APINBNO3 (the 5th most abundant species in the aerosol phase) where it is an order of magnitude lower than the next highest estimate. Therefore, systematic underestimation of vapor pressure is not the obvious cause of overabundance of this product in our simulation.

3.2.2 Model-measurement comparison of FG mole fractions

Comparing with observations, discrepancies in the proportions of $CONO_2$ and naCO are higher than in the APIN-INO_x case. $CONO_2$ mole fraction is overestimated by the model as it accounts for 6% of the relative mole fraction after 20 hours, while in the measurements it accounts for only 2% of the relative mole fraction (Figure 3). The model also over predicts the relative mole fraction of naCO (7% against less than 1% in the measurements).

The low relative humidity conditions of the experiments (RH <5%) exclude organonitrate hydrolysis (Liu et al., 2012), not included in the model, as a possible condensed phase pathway that explains the model-measurement discrepancy for CONO₂. Organonitrate compounds are formed from the addition of NO to a peroxy radical (e.g., C813NO3 is formed from the addition of NO to C813O2). Yields are affected by the rate of HO₂ or NO₃ addition to the peroxy radical, and the branching ratio of the reaction to produce organonitrate or alkoxy radical and NO₂ (Noziere et al., 1999; Ruppert et al., 1999; Aschmann et al., 2002; Pinho et al., 2007). High uncertainty in CONO₂ production rates by lumped chemical reaction schemes has also been reported (Henderson et al., 2011), but uncertainties may also be present in explicit mechanisms for the reasons described. Smaller number of components condensing to the aerosol phase may lead to greater sensitivity of simulation results to individual values of such rate constants or vapor pressures (which may otherwise be compensated across a larger suite of compounds or reactions), resulting in higher likelihood of discrepancies between model predictions and observations.

3.3 APIN-nNO_x

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3.3.1 Simulation results

The apportionment of O:C in the APIN- nNO_x system is qualitatively similar to APIN- $1NO_x$, sans contributions from nitrogenated groups. The FG composition of gas and aerosol O:C ratios are very similar, though the value is higher in the latter phase (Figure 1). The aerosol-phase O:C ratio increases in the simulation to arrive to 0.53 after 20 hours, while the observed O:C ratio by Chhabra et al. (2011) is between 0.3–0.4. We can see that in the very beginning of the simulations the only compounds contributing to the O:C ratio that are able to partition to the aerosol phase have aCOH and hydroperoxide moieties. The ketone FG starts contributing to O:C in the aerosol phase only after this initial phase. The hydroperoxide FG accounts for 42% of the total SOA O:C ratio after 20 hours of simulation.

As for APIN-hNO_x, the $C_{\rm SOA}$ formed in these simulations is an order of magnitude less than the 64 $\mu \rm g \, m^{-3}$ reported for the corresponding experiment. Sensitivity analysis with respect to $C_{\rm SOA}$ (by varying the amount of absorptive mass) indicates that relative proportions analyzed are again representative for aerosol formed according to the chemical mechanism and vapor pressure estimation method in our simulation framework (Appendix B), which excludes accretion reactions in the condensed phase. The same four carbonyl compounds that make up 70% of the ketone CO comprise over 90% of the naCO, 80% of hydroperoxide, and 80% of the $C_{\rm SOA}$ in this system (Figure 7) at peak $C_{\rm SOA}$ (12.1 hours). In contrast to APIN-lNO_x and APIN-hNO_x, multifunctional organonitrate compounds do not contribute to the COOH abundance; it is effectively accounted for by only two compounds: H3C25CCO2H (Table 2) contributing 60% and PINONIC (pinonic acid) contributing 40%. The reason for the large contribution of H3C25CCO2H is its saturation concentration of log C⁰ of 1.2, while PINONIC has a log C⁰ of

3.0 but its total (gas+aerosol) concentration is almost an order of magnitude more than H3C25CCO2H (Figure S13). As with the APIN-hNO_x simulations, the dominance of so few polyfunctional compounds in the aerosol phase is surprising; past studies have identified more than five smaller compounds comprising observed APIN (and other precursor) ozonolysis aerosol yields under dry conditions (Yu et al., 1999; Pankow, 2001). These compounds are primarily composed of COOH, aCOH, and aldehyde CO groups, which are present in low abundance in our simulations.

3.3.2 Model-measurement comparison of FG mole fractions

In Figure 3 we can observe that the highest discrepancies in the FG relative mole fraction between experimental observations and simulations are found in the oxygenated FGs (COOH, aCOH, and naCO). While the naCO mass fraction is overestimated by the model (7% in the model against 3% in the experiment), the COOH and the aCOH are underestimated (less than 1% in the model against 3% in the experiment for COOH and 7% against 15% for aCOH). Reactions of aldehydes with hydroperoxides can form peroxyhemiacetals (Jang and Kamens, 2001; Docherty et al., 2005), leading to a reduction in naCO. However, if the condensed-phase naCO is mostly ketone as predicted by the model (Figures 1 and S8), this is not likely to improve model-measurement agreement of the relative mole fractions of naCO.

Simulation of COOH production by gas-phase oxidation has also been reported to underestimate its abundance in OA in other studies (e.g., Aumont et al., 2012). In particular, there is a question whether the gas-phase production rate is low or production and degradation rates are both high. To examine this question, IRR contributions to production and loss of COOH from semi-volatile compounds that can condense to the aerosol phase in appreciable proportions ($\log C^0 \le 2.5$) are shown in Figure 8. The net change in moles of COOH due to degradation is 83% of the production for these compounds. There are more compounds contributing to the aerosol-phase COOH in the APIN-lNO_x simulation and the net degradation is only 13% of the net production. One known mechanism for production of COOH by heterogeneous reactions not included in the model involves the transformation of hydroxy carbonyls (formed from alkoxy radicals) to dihydrofurans (Ziemann and Atkinson, 2012), which are further oxidized in the gas phase by O_3 , primarily. However, as there is no NO_x in the APIN-nNO_x system, the O_3 production rate is small (concentrations are less than 0.6 ppb in our simulations, Figure S14). The model-measurement discrepancy may again be partially due to the low concentrations of condensible products and small number of products partitioning to the aerosol phase in this simulation. Production and condensation of ELVOCs or additional oxidation mechanisms in the condensed phase not implemented in our model may also contribute to this discrepancy.

3.4 TMB-INO_{x}

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3.4.1 Simulation results

In the TMB-INO_x simulations, continued oxidation in the gas-phase proceeds for the entire duration of simulation and the O:C ratio approaches unity, while the aerosol-phase O:C and FG composition largely stabilizes in magnitude after the first several hours. In this mechanism, we note the presence of esters, ethers, organic peroxides (ROOR' + ROOH), and anhydrides which were not present in the APIN photooxidation schemes. The O:C ratio of the simulated SOA is (\sim 0.7). While overall O:C in

experiments of Sax et al. (2005) were not reported, Sato et al. (2012) report values in the range of 0.25–0.47 for similar TMB photooxidation experiments with NO_x and methyl nitrate as OH source, with no seed and lower RH (<1%).

The peroxyacyl nitrate is the FG that contributes the most to the gas phase O:C ratio (around 30%), as in the APIN-INO_x case. Peroxide, hydroperoxide, and CONO₂ are major contributors to the aerosol O:C ratio and ester and aCOH are also present in the aerosol phase. COOH, anhydrides, peroxy acid, peroxyacyl nitrate and naCO contribute to the gas phase. The high peroxide and hydroperoxide contribution to the aerosol phase in this simulation agrees with their reported role in SOA formation from TMB photooxidation in low NO_x conditions (Wyche et al., 2009). However, the C_{SOA} mass fraction of compounds containing organic peroxide is 96% in our simulations, which is higher than what has been experimentally determined by Sato et al. (2012) in similar conditions (12±8%).

While higher precursor concentrations lead to high concentrations of condensible products in the gas-phase, two compounds (TM135BPOOH and NMXYFUOOH) make up over 70% of $C_{\rm SOA}$ at its peak (11.9 hours). While formation and condensation of ELVOCs in the experimental system cannot be ruled out, it is likely that their contribution would be much smaller than the SVOC fraction on account of the high mass loadings (Figure S1) (Zhang et al., 2015). TM135BPOOH is found to contribute 60% to the aCH and 65% of hydroperoxide FGs in the aerosol phase at peak $C_{\rm SOA}$ (Figure 9). TM135BPOOH is an hydroxy hydroperoxide bicyclic peroxide (Table 2) that is formed after many oxidation steps that follow the first addition of OH to the aromatic ring (Rickard et al., 2010). This compound has been also indicated by Rickard et al. (2010) as a potential SOA forming compound from TMB photooxidation and was found to be the most abundant hydroperoxide compound in the beginning of the photooxidation simulation they conducted. In our estimation, vapor pressures for these compounds predicted by SIMPOL.1 are at the upper end of estimates by three other methods (Table 3) so it is ress plausible that the high abundance of these species in the condensed phase can be attributed to understimation of vapor pressures.

3.4.2 Model-measurement comparison of FG mole fractions

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The model is able to capture the general trends in COOH and $CONO_2$ relative to the first sample but the difference in magnitude is higher than APIN-lNO_x case (Figure 2). The measured changes in COOH indicate an increase of a factor of four, while predicted COOH increases by as much as a factor of eight during the same time. The $CONO_2$ mole fraction at the end of our simulation is around 0.7, while it arrives to 0.4 in the experimental observations. The increase in naCO and aCOH that are captured by simulation for APIN-lNO_x are not captured for TMB-lNO_x. As for APIN-lNO_x, the degradation of the aCH fraction and the appearance of naCO and aCOH are slower in the model than in the experiments carried by Sax et al. (2005). From the molar fractions shown at different simulation times in Figure 3 we find better agreement in the beginning of the simulation and the differences in the FG molar fraction is higher compared to APIN-lNO_x.

The low proportion of naCO compared to simulations described in preceding sections is due to the lack of multifunctional ketone compounds. The ester CO abundance is on the same order as ketone CO (Figure S8) in our simulations, and it is possible that the naCO reported by Sax et al. (2005) may include ester CO as the absorption band at \sim 1735 cm $^{-1}$ is close to aldehyde and ketone CO at \sim 1725 and \sim 1715 cm $^{-1}$, respectively (Pavia et al., 2008). However, the ester CO contribution cannot explain the entire difference given the large discrepancy. Peroxide and hydroperoxide photolysis in the condensed phase

under UV irradiation can lead to the increase of both naCO and aCOH FGs, but increases calculated using the 6-day lifetime (Epstein et al., 2014) only partially explain this difference (with an increase from 7 to 8% for aCOH and from 1 to 3% for CO after 20 hours of irradiation). As there is indication that O:C and abundance of hydroperoxide and peroxide groups may be overestimated, it is possible that over-representation of TM135BPOOH in the simulated aerosol phase (Section 3.4.1) also contributes to an overestimation of aCH, leading to a smaller fraction of the measured oxygenated groups (COOH, aCOH, and naCO).

4 Conclusions

In this study, the FG distribution of SOA generated in environmental control chamber experiments reported in literature for APIN and TMB photooxidation have been compared to explicit gas-phase chemistry and partitioning simulated with MCMv3.2 and SIMPOL.1. Varying degrees of agreement between the model and FTIR measurements of FG evolution in SOA generated in environmental controlled chambers are found.

In the APIN-INO_x simulations, the FG relative abundance is well captured by the model in the first hours of simulation, and general trends in the changes of the mole fraction compared to the first sample are captured qualitatively by the model. However, the underestimation of the measured oxidized groups (COOH, aCOH, and CO) are apparent after 20 hours in our simulations; this discrepancy may be explained by heterogeneous reactions missing in the model. O:C is generally overestimated for APIN-hNO_x, APIN-nNO_x, and TMB-lNO_x on account of large contributions from CONO₂, peroxide, or hydroperoxide groups, while the aCH is simulated consistently in larger proportion to some of the measured oxygenated species (COOH and aCOH). These errors are largely correlated, as C_{SOA} mass and individual FGs are dominated by a few polyfunctional compounds in these simulations. The dependencies of aerosol composition on a limited number of compounds also speaks to the sensitivity of simulation results on a few kinetic or partitioning parameters, which might otherwise be averaged out in systems where the condensed phase is composed of a larger number of compounds.

In the APIN-hNO_x simulations, the model predicts a higher fractional abundance of CONO₂ in the aerosol phase than what is observed in the FTIR measurements. The CONO₂ fraction arrives to constitute 46% of the total O:C ratio, which partly contributes to the higher O:C of the aerosol phase during the simulation (0.78) compared to the O:C observed (\sim 0.4) by Chhabra et al. (2011). Only four CONO₂-containing polyfunctional compounds account for more than 80% of the organic mass. The uncertainties due to lack of kinetic data in the total CONO₂ yield in the primary oxidation sequence of APIN may play an important role in the high NO_x regime and explain the discrepancies between model and measurements in this scenario. For the APIN-nNO_x simulations, four polyfunctional compounds account for over 80% of the C_{SOA} mass and a large bulk of ketone CO and hydroperoxide FGs. The relative abundance of ketone CO is overestimated compared to observations; the O:C is also overestimated, possibly on account of the large (42%) contribution from the hydroperoxide FG which originates from the same set of molecules. For the TMB-lNO_x photooxidation simulations, general trends in the changes in relative mole fractions compared to the first sample for COOH, aCOH, naCO, aCH and CONO₂ also qualitatively follow observations, but their magnitudes have more discrepancies with experiments than in the case of APIN-lNO_x. These discrepancies have also

been hypothesized as a sensitivity to reaction rates and vapor pressures of a few dominant products that contribute significantly to the aCH mole fraction and peroxide fraction of the aerosol O:C ratio. As for the APIN-lNO $_x$ simulations, the agreement in abundances of aCH relative to the measured set of oxidized FGs may also be explained by additional condensed phase oxidation chemistry not included in the model.

This work illustrates that concurrent measurement of FGs alongside common techniques for atomic and molecular characterization of OA can provide an opportunity for complementary evaluation and further guide detailed understanding of chemical and physical transformations. Analysis of FG abundance can supplement tracking of individual tracers and evaluate the importance of mechanisms that lead to production of a class of compounds in the overall molar (or mass) budget. FG abundances can also provide structural interpretation to variations in elemental ratios (e.g., O:C, H:C, and N:C). Looking forward, systematic model-measurement comparison of FGs under controlled conditions may be able to provide constraints and aid development of chemical mechanism generators (e.g., Gao et al., 2016; Aumont et al., 2005). While we have uncovered only a fraction of the analysis capabilities that a FG perspective provides, we anticipate that the tools and approaches introduced in this work can encourage further comparisons between model simulations of both gas and aerosol-phase chemistry in conjunction with emerging methods for FG quantification.

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References

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- Aimanant, S. and Ziemann, P. J.: Development of Spectrophotometric Methods for the Analysis of Functional Groups in Oxidized Organic Aerosol, Aerosol Science and Technology, 47, 581–591, doi:10.1080/02786826.2013.773579, 2013.
- Allen, D. T., Palen, E. J., Haimov, M. I., Hering, S. V., and Young, J. R.: Fourier-transform Infrared-spectroscopy of Aerosol Collected
 In A Low-pressure Impactor (LPI/FTIR) Method Development and Field Calibration, Aerosol Science and Technology, 21, 325–342, doi:10.1080/02786829408959719, 1994.
 - Aschmann, S. M., Atkinson, R., and Arey, J.: Products of reaction of OH radicals with α -pinene, Journal of Geophysical Research: Atmospheres, 107, ACH 6–1–ACH 6–7, doi:10.1029/2001JD001098, 2002.
- Aumont, B., Szopa, S., and Madronich, S.: Modelling the evolution of organic carbon during its gas-phase tropospheric oxidation: development of an explicit model based on a self generating approach, Atmospheric Chemistry and Physics, 5, 2497–2517, doi:10.5194/acp-5-2497-2005, 2005.
 - Aumont, B., Valorso, R., Mouchel-Vallon, C., Camredon, M., Lee-Taylor, J., and Madronich, S.: Modeling SOA formation from the oxidation of intermediate volatility n-alkanes, Atmospheric Chemistry and Physics, 12, 7577–7589, doi:10.5194/acp-12-7577-2012, 2012.
 - Barley, M. H., Topping, D., Lowe, D., Utembe, S., and McFiggans, G.: The sensitivity of secondary organic aerosol (SOA) component partitioning to the predictions of component properties Part 3: Investigation of condensed compounds generated by a near-explicit model of VOC oxidation, Atmospheric Chemistry and Physics, 11, 13 145–13 159, doi:10.5194/acp-11-13145-2011, 2011.
 - Bloss, C., Wagner, V., Jenkin, M. E., Volkamer, R., Bloss, W. J., Lee, J. D., Heard, D. E., Wirtz, K., Martin-Reviejo, M., Rea, G., Wenger, J. C., and Pilling, M. J.: Development of a detailed chemical mechanism (MCMv3.1) for the atmospheric oxidation of aromatic hydrocarbons, Atmospheric Chemistry and Physics, 5, 641–664, doi:10.5194/acp-5-641-2005, 2005.
- Bonn, B., von Kuhlmann, R., and Lawrence, M. G.: High contribution of biogenic hydroperoxides to secondary organic aerosol formation, Geophysical Research Letters, 31, L10 108, doi:10.1029/2003GL019172, 2004.
 - Capouet, M., Mueller, J. F., Ceulemans, K., Compernolle, S., Vereecken, L., and Peeters, J.: Modeling aerosol formation in alpha-pinene photo-oxidation experiments, Journal of Geophysical Research-atmospheres, 113, D02 308, doi:10.1029/2007JD008995, 2008.
 - Cappa, C. D.: A model of aerosol evaporation kinetics in a thermodenuder, Atmospheric Measurement Techniques, 3, 579–592, doi:10.5194/amt-3-579-2010, 2010.
 - Cappa, C. D. and Wilson, K. R.: Multi-generation gas-phase oxidation, equilibrium partitioning, and the formation and evolution of secondary organic aerosol, Atmospheric Chemistry and Physics, 12, 9505–9528, doi:10.5194/acp-12-9505-2012, 2012.
 - Cappa, C. D., Jathar, S. H., Kleeman, M. J., Docherty, K. S., Jimenez, J. L., Seinfeld, J. H., and Wexler, A. S.: Simulating secondary organic aerosol in a regional air quality model using the statistical oxidation model Part 2: Assessing the influence of vapor wall losses, Atmospheric Chemistry and Physics, 16, 3041–3059, doi:10.5194/acp-16-3041-2016, 2016.
 - Chameides, W., Lindsay, R., Richardson, J., and Kiang, C.: The role of biogenic hydrocarbons in urban photochemical smog: Atlanta as a case study, Science, 241, 1473–1475, doi:10.1126/science.3420404, 1988.
 - Chang, E. I. and Pankow, J. F.: Organic particulate matter formation at varying relative humidity using surrogate secondary and primary organic compounds with activity corrections in the condensed phase obtained using a method based on the Wilson equation, Atmospheric Chemistry and Physics, 10, 5475–5490, doi:10.5194/acp-10-5475-2010, 2010.

- Chen, Q., Liu, Y., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-phase chemistry of secondary organic material: modeled compared to measured O:C and H:C elemental ratios provide constraints., Environmental science & technology, 45, 4763–70, doi:10.1021/es104398s, 2011.
- Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan, R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic aerosol, Atmospheric Chemistry and Physics, 11, 8827–8845, doi:10.5194/acp-11-8827-2011, 2011.
 - Cleveland, M. J., Ziemba, L. D., Griffin, R. J., Dibb, J. E., Anderson, C. H., Lefer, B., and Rappenglück, B.: Characterization of urban aerosol using aerosol mass spectrometry and proton nuclear magnetic resonance spectroscopy, Atmospheric Environment, 54, 511 518, doi:10.1016/j.atmosenv.2012.02.074, 2012.
- 10 Cocker, D. R., Clegg, S. L., Flagan, R. C., and Seinfeld, J. H.: The effect of water on gas-particle partitioning of secondary organic aerosol. Part I: alpha-pinene/ozone system, Atmospheric Environment, 35, 6049–6072, doi:10.1016/S1352-2310(01)00404-6, 2001.
 - Compernolle, S., Ceulemans, K., and Muller, J. . F.: EVAPORATION: a new vapour pressure estimation methodfor organic molecules including non-additivity and intramolecular interactions, Atmospheric Chemistry and Physics, 11, 9431–9450, doi:10.5194/acp-11-9431-2011, 2011.
- 15 Coury, C. and Dillner, A. M.: A method to quantify organic functional groups and inorganic compounds in ambient aerosols using attenuated total reflectance FTIR spectroscopy and multivariate chemometric techniques, Atmospheric Environment, 42, 5923–5932, doi:10.1016/i.atmosenv.2008.03.026, 2008.
 - Damian, V., Sandu, A., Damian, M., Potra, F., and Carmichael, G. R.: The kinetic preprocessor KPP-a software environment for solving chemical kinetics, Computers & Chemical Engineering, 26, 1567 1579, doi:http://dx.doi.org/10.1016/S0098-1354(02)00128-X, 2002.
- Day, D. A., Liu, S., Russell, L. M., and Ziemann, P. J.: Organonitrate group concentrations in submicron particles with high nitrate and organic fractions in coastal southern California, Atmospheric Environment, 44, 1970–1979, doi:10.1016/j.atmosenv.2010.02.045, 2010.
 - DAYLIGHT Chemical Information Systems, I.: SMARTS A Language for Describing Molecular Patterns, http://www.daylight.com/dayhtml/doc/theory/smarts.html, accessed 30 September 2015, Accessed 30 September 2015.
- Decesari, S., Mircea, M., Cavalli, F., Fuzzi, S., Moretti, F., Tagliavini, E., and Facchini, M. C.: Source attribution of water-soluble organic aerosol by nuclear magnetic resonance spectroscopy, Environmental Science & Technology, 41, 2479–2484, doi:10.1021/es0617111, 2007.
 - Denbigh, K.: The Principles of Chemical Equilibrium, Cambridge University Press, Cambridge, UK, 1981.

- Derwent, R., Jenkin, M., and Saunders, S.: Photochemical ozone creation potentials for a large number of reactive hydrocarbons under European conditions, Atmospheric Environment, 30, 181–199, doi:10.1016/1352-2310(95)00303-G, 1996.
- Derwent, R. G., Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: Photochemical ozone creation potentials for organic compounds in northwest Europe calculated with a master chemical mechanism, Atmospheric Environment, 32, 2429–2441, doi:10.1016/S1352-2310(98)00053-3, 1998.
 - Docherty, K. S., Wu, W., Lim, Y. B., and Ziemann, P. J.: Contributions of organic peroxides to secondary aerosol formed from reactions of monoterpenes with O-3, Environmental Science & Technology, 39, 4049–4059, doi:10.1021/es050228s, 2005.
 - Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile organics, Environmental Science & Technology, 40, 2635–2643, doi:10.1021/es052297c, 2006.
 - Donahue, N. M., Kroll, J. H., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set Part 2: Diagnostics of organicaerosol evolution, Atmospheric Chemistry and Physics, 12, 615–634, doi:10.5194/acp-12-615-2012, 2012.

- Donahue, N. M., Chuang, W., Epstein, S. A., Kroll, J. H., Worsnop, D. R., Robinson, A. L., Adams, P. J., and Pandis, S. N.: Why do organic aerosols exist? Understanding aerosol lifetimes using the two-dimensional volatility basis set, Environmental Chemistry, 10, 151, doi:10.1071/EN13022, 2013.
- Dron, J., El Haddad, I., Temime-Roussel, B., Jaffrezo, J.-L., Wortham, H., and Marchand, N.: Functional group composition of ambient and source organic aerosols determined by tandem mass spectrometry, Atmospheric Chemistry and Physics, 10, 7041–7055, doi:10.5194/acp-10-7041-2010, 2010.
 - Eddingsaas, N. C., Loza, C. L., Yee, L. D., Seinfeld, J. H., and Wennberg, P. O.: alpha-pinene photooxidation under controlled chemical conditions Part 1: Gas-phase composition in low- and high-NOx environments, Atmospheric Chemistry and Physics, 12, 6489–6504, doi:10.5194/acp-12-6489-2012, 2012.
- 10 Ehn, M., Thornton, J. A., Kleist, E., Sipilía, M., Junninen, H., Pullinen, I., Springer, M., Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I.-H., Rissanen, M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T., Kurtén, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Maso, M. D., Berndt, T., Petíajía, T., Wahner, A., Kerminen, V.-M., Kulmala, M., Worsnop, D. R., Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol, Nature, 506, 476–479, doi:10.1038/nature13032, 2014.
- 15 Epstein, S. A. and Donahue, N. M.: Ozonolysis of Cyclic Alkenes as Surrogates for Biogenic Terpenes: Primary Ozonide Formation and Decomposition, Journal of Physical Chemistry A, 114, 7509–7515, doi:10.1021/jp102177v, 2010.
 - Epstein, S. A., Blair, S. L., and Nizkorodov, S. A.: Direct Photolysis of a-Pinene Ozonolysis Secondary Organic Aerosol: Effect on Particle Mass and Peroxide Content, Environmental Science & Technology, 48, 11 251–11 258, doi:10.1021/es502350u, 2014.
 - Fan, Y., Qin, F., Luo, X., Lin, L., Gui, H., and Liu, J.: Heterogeneous condensation on insoluble spherical particles: Modeling and parametric study, Chemical Engineering Science, 102, 387 396, doi:10.1016/j.ces.2013.08.040, 2013.

- Fehsenfeld, F., Calvert, J., Fall, R., Goldan, P., Guenther, A. B., Hewitt, C. N., Lamb, B., Liu, S., Trainer, M., Westberg, H., and Zimmerman, P.: Emissions of volatile organic compounds from vegetation and the implications for atmospheric chemistry, Global Biogeochemical Cycles, 6, 389–430, doi:10.1029/92GB02125, 1992.
- Flagan, R. C. and Seinfeld, J. H.: Fundamentals of air pollution engineering, Prentice-Hall, Inc., Englewood Cliffs, New Jersey, 1988.
- Fuchs, N. A. and Sutugin, A. G.: High-Dispersed {Aerosols}, in: Topics in Current Aerosol Research, edited by Brock, G. M. and Hidy, J. R., International Reviews in Aerosol Physics and Chemistry, pp. 1–3, Pergamon, doi:10.1016/B978-0-08-016674-2.50006-6, 1971.
 - Fuentes, J. D., Gu, L., Lerdau, M., Atkinson, R., Baldocchi, D., Bottenheim, J. W., Ciccioli, P., Lamb, B., Geron, C., Guenther, A., Sharkey, T. D., and Stockwell, W.: Biogenic Hydrocarbons in the Atmospheric Boundary Layer: A Review, Bull. Amer. Meteor. Soc., 81, 1537–1575, doi:10.1175/1520-0477(2000)081<1537:BHITAB>2.3.CO;2, 2000.
- Gao, C. W., Allen, J. W., Green, W. H., and West, R. H.: Reaction Mechanism Generator: Automatic construction of chemical kinetic mechanisms, Computer Physics Communications, 203, 212 225, doi:http://dx.doi.org/10.1016/j.cpc.2016.02.013, 2016.
 - Gao, S., Keywood, M., Ng, N. L., Surratt, J., Varutbangkul, V., Bahreini, R., Flagan, R. C., and Seinfeld, J. H.: Low-molecular-weight and oligomeric components in secondary organic aerosol from the ozonolysis of cycloalkenes and alpha-pinene, Journal of Physical Chemistry A, 108, 10147–10164, doi:10.1021/jp047466e, 2004.
- 35 Grieshop, A. P., Miracolo, M. A., Donahue, N. M., and Robinson, A. L.: Constraining the Volatility Distribution and Gas-Particle Partitioning of Combustion Aerosols Using Isothermal Dilution and Thermodenuder Measurements, Environmental Science & Technology, 43, 4750– 4756, doi:10.1021/es8032378, 2009.

- Griffin, R. J., Nguyen, K., Dabdub, D., and Seinfeld, J. H.: A coupled hydrophobic-hydrophilic model for predicting secondary organic aerosol formation, Journal of Atmospheric Chemistry, 44, 171–190, doi:10.1023/A:1022436813699, 2003.
- Hall, W. A. and Johnston, M. V.: Oligomer Content of alpha-Pinene Secondary Organic Aerosol, Aerosol Science and Technology, 45, 37–45, doi:10.1080/02786826.2010.517580, 2011.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and emerging issues, Atmospheric Chemistry and Physics, 9, 5155–5236, doi:10.5194/acp-9-5155-2009, 2009.
- Hamilton, J. F., Webb, P. J., Lewis, A. C., Hopkins, J. R., Smith, S., and Davy, P.: Partially oxidised organic components in urban aerosol using GCXGC-TOF/MS, Atmospheric Chemistry and Physics, 4, 1279–1290, doi:doi:10.5194/acp-4-1279-2004, 2004.
 - Hayman, G. D.: Effects of Pollution Control on UV Exposure, AEA Technology Final Report (Reference AEA/RCEC/22522001/R/002 ISSUE1) prepared for the Department of Health on Contract 121/6377, Tech. rep., AEA Technology, Oxfordshire, UK, 1997.
- Hemming, B. L. and Seinfeld, J. H.: On the Hygroscopic Behavior of Atmospheric Organic Aerosols, Industrial & Engineering Chemistry

 Research, 40, 4162–4171, doi:10.1021/ie000790l, 2001.
 - Henderson, B. H.: Python-based Environment for Reaction Mechanisms/Mathematics (PERMM), doi:dx.doi.org/10.5281/zenodo.44396, https://github.com/barronh/permm/, 2015.
 - Henderson, B. H.: Kinetic Pre-Processor with updates to allow working with MCM, doi:10.5281/zenodo.44682, http://github.com/barronh/kpp, 2016.
- 20 Henderson, B. H., Pinder, R. W., Crooks, J., Cohen, R. C., Hutzell, W. T., Sarwar, G., Goliff, W. S., Stockwell, W. R., Fahr, A., Mathur, R., Carlton, A. G., and Vizuete, W.: Evaluation of simulated photochemical partitioning of oxidized nitrogen in the upper troposphere, Atmospheric Chemistry and Physics, 11, 275–291, doi:10.5194/acp-11-275-2011, 2011.
 - Holes, A., Eusebi, A., Grosjean, D., and Allen, D. T.: FTIR analysis of aerosol formed in the photooxidation of 1,3,5-trimethylbenzene, Aerosol Science and Technology, 26, 516–526, doi:10.1080/02786829708965450, 1997.
- 25 Im, Y., Jang, M., and Beardsley, R. L.: Simulation of aromatic SOA formation using the lumping model integrated with explicit gas-phase kinetic mechanisms and aerosol-phase reactions, Atmospheric Chemistry and Physics, 14, 4013–4027, doi:10.5194/acp-14-4013-2014, 2014.
 - Jang, M. S. and Kamens, R. M.: Characterization of secondary aerosol from the photooxidation of toluene in the presence of NOx and 1-propene, Environmental Science & Technology, 35, 3626–3639, doi:10.1021/es010676+, 2001.
- Jathar, S. H., Cappa, C. D., Wexler, A. S., Seinfeld, J. H., and Kleeman, M. J.: Multi-generational oxidation model to simulate secondary organic aerosol in a 3-D air quality model, Geoscientific Model Development, 8, 2553–2567, doi:10.5194/gmd-8-2553-2015, 2015.
 - Jayne, J. T., Leard, D. C., Zhang, X. F., Davidovits, P., Smith, K. A., Kolb, C. E., and Worsnop, D. R.: Development of an aerosol mass spectrometer for size and composition analysis of submicron particles, Aerosol Science and Technology, 33, 49–70, doi:10.1080/027868200410840, 2000.
- Jeffries, H. E. and Tonnesen, S.: A comparison of two photochemical reaction mechanisms using mass balance and process analysis, Atmospheric Environment, 28, 2991 3003, doi:10.1016/1352-2310(94)90345-X, 1994.
 - Jenkin, M. E.: Modelling the formation and composition of secondary organic aerosol from alpha- and beta-pinene ozonolysis using MCM v3, Atmospheric Chemistry and Physics, 4, 1741–1757, doi:10.5194/acp-4-1741-2004, 2004.

- Jenkin, M. E., Saunders, S. M., and Pilling, M. J.: The tropospheric degradation of volatile organic compounds: a protocol for mechanism development, Atmospheric Environment, 31, 81 104, doi:10.1016/S1352-2310(96)00105-7, 1997.
- Jenkin, M. E., Saunders, S. M., Wagner, V., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part B): tropospheric degradation of aromatic volatile organic compounds, Atmospheric Chemistry and Physics, 3, 181–193, doi:10.5194/acp-3-181-2003, 2003.

20

25

- Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, W. G. L., L. D., G. R., Williams, P. I., Bower, K., Kondo, W. G. L., L. D., G. R., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, W. G. L., L. D., G. R., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., G. R., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Collins, D. R., Cubison, M. J., Dunlea, E. J., Williams, P. J., Charles, M. R., Williams, P
- Y., Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the Atmosphere, Science, 326, 1525–1529, doi:10.1126/science.1180353, 2009.
- Kalberer, M., Paulsen, D., Sax, M., Steinbacher, M., Dommen, J., Prevot, A. S. H., Fisseha, R., Weingartner, E., Frankevich, V., Zenobi,
 R., and Baltensperger, U.: Identification of polymers as major components of atmospheric organic aerosols, Science, 303, 1659–1662,
 doi:10.1126/science.1092185, 2004.
 - Kalberer, M., Sax, M., and Samburova, V.: Molecular size evolution of oligomers in organic aerosols collected in urban atmospheres and generated in a smog chamber, Environmental Science & Technology, 40, 5917–5922, doi:10.1021/es0525760, 2006.
 - Kostenidou, E., Pathak, R. K., and Pandis, S. N.: An algorithm for the calculation of secondary organic aerosol density combining AMS and SMPS data, Aerosol Science and Technology, 41, 1002–1010, doi:10.1080/02786820701666270, 2007.
 - Kroll, J. H. and Seinfeld, J. H.: Chemistry of secondary organic aerosol: Formation and evolution of low-volatility organics in the atmosphere, Atmospheric Environment, 42, 3593–3624, doi:10.1016/j.atmosenv.2008.01.003, 2008.
 - Kroll, J. H., Donahue, N. M., Jimenez, J. L., Kessler, S. H., Canagaratna, M. R., Wilson, K. R., Altieri, K. E., Mazzoleni, L. R., Wozniak, A. S., Bluhm, H., Mysak, E. R., Smith, J. D., Kolb, C. E., and Worsnop, D. R.: Carbon oxidation state as a metric for describing the chemistry of atmospheric organic aerosol, Nature Chemistry, 3, 133–139, doi:10.1038/nchem.948, 2011.
 - La, Y. S., Camredon, M., Ziemann, P. J., Valorso, R., Matsunaga, A., Lannuque, V., Lee-Taylor, J., Hodzic, A., Madronich, S., and Aumont, B.: Impact of chamber wall loss of gaseous organic compounds on secondary organic aerosol formation: explicit modeling of SOA formation from alkane and alkene oxidation, Atmospheric Chemistry and Physics, 16, 1417–1431, doi:10.5194/acp-16-1417-2016, 2016.
 - Lamb, B., Gay, D., Westberg, H., and Pierce, T.: A biogenic hydrocarbon emission inventory for the U.S.A. using a simple forest canopy model, Atmospheric Environment. Part A. General Topics, 27, 1673 1690, doi:10.1016/0960-1686(93)90230-V, 1993.
 - Lee, A. K. and Chan, C. K.: Single particle Raman spectroscopy for investigating atmospheric heterogeneous reactions of organic aerosols, Atmospheric Environment, 41, 4611 4621, doi:http://dx.doi.org/10.1016/j.atmosenv.2007.03.040, 2007.
 - Liu, S., Shilling, J. E., Song, C., Hiranuma, N., Zaveri, R. A., and Russell, L. M.: Hydrolysis of Organonitrate Functional Groups in Aerosol Particles, Aerosol Science and Technology, 46, 1359–1369, doi:10.1080/02786826.2012.716175, 2012.
- Loza, C. L., Chan, A. W. H., Galloway, M. M., Keutsch, F. N., Flagan, R. C., and Seinfeld, J. H.: Characterization of Vapor Wall Loss in Laboratory Chambers, Environmental Science & Technology, 44, 5074–5078, doi:10.1021/es100727v, 2010.
 - Matsunaga, A. and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield Measurements, Aerosol Science and Technology, 44, 881–892, doi:10.1080/02786826.2010.501044, 2010.

- McFiggans, G., Topping, D. O., and Barley, M. H.: The sensitivity of secondary organic aerosol component partitioning to the predictions of component properties Part 1: A systematic evaluation of some available estimation techniques, Atmospheric Chemistry and Physics, 10, 10 255–10 272, doi:10.5194/acp-10-10255-2010, 2010.
- McVay, R. C., Zhang, X., Aumont, B., Valorso, R., Camredon, M., La, Y. S., Wennberg, P. O., and Seinfeld, J. H.: SOA formation from the photooxidation of α-pinene: systematic exploration of the simulation of chamber data, Atmospheric Chemistry and Physics, 16, 2785–2802, doi:10.5194/acp-16-2785-2016, 2016.

- Mertes, P., Pfaffenberger, L., Dommen, J., Kalberer, M., and Baltensperger, U.: Development of a sensitive long path absorption photometer to quantify peroxides in aerosol particles (Peroxide-LOPAP), Atmospheric Measurement Techniques, 5, 2339–2348, doi:10.5194/amt-5-2339-2012, 2012.
- 10 Metzger, A., Dommen, J., Gaeggeler, K., Duplissy, J., Prevot, A. S. H., Kleffmann, J., Elshorbany, Y., Wisthaler, A., and Baltensperger, U.: Evaluation of 1,3,5 trimethylbenzene degradation in the detailed tropospheric chemistry mechanism, MCMv3.1, using environmental chamber data, Atmospheric Chemistry and Physics, 8, 6453–6468, doi:10.5194/acp-8-6453-2008, 2008.
 - Ming, Y. and Russell, L. M.: Thermodynamic equilibrium of organic-electrolyte mixtures in aerosol particles, Aiche Journal, 48, 1331–1348, 2002.
- Mozurkewich, M. and Benson, S. W.: Self-Reaction of HO2 and DO2: Negative temperature dependence and pressure effects, International Journal of Chemical Kinetics, 17, 787–807, doi:10.1002/kin.550170802, 1985.
 - Murphy, B. N., Donahue, N. M., Fountoukis, C., and Pandis, S. N.: Simulating the oxygen content of ambient organic aerosol with the 2D volatility basis set, Atmospheric Chemistry and Physics, 11, 7859–7873, doi:10.5194/acp-11-7859-2011, 2011.
 - Myrdal, P. B. and Yalkowsky, S. H.: Estimating pure component vapor pressures of complex organic molecules, Industrial & Engineering Chemistry Research, 36, 2494–2499, doi:10.1021/ie950242l, 1997.
 - Nannoolal, Y., Rarey, J., and Ramjugernath, D.: Estimation of pure component properties Part 3. Estimation of the vapor pressure of non-electrolyte organic compounds via group contributions and group interactions, Fluid Phase Equilibria, 269, 117–133, doi:10.1016/j.fluid.2008.04.020, 2008.
- Ng, N. L., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Kroll, J. H., Kwan, A. J., McCabe, D. C., Wennberg, P. O., Sorooshian, A., Murphy,
 S. M., Dalleska, N. F., Flagan, R. C., and Seinfeld, J. H.: Effect of NO(x) level on secondary organic aerosol (SOA) formation from the photooxidation of terpenes, Atmospheric Chemistry and Physics, 7, 5159–5174, doi:10.5194/acp-7-5159-2007, 2007.
 - Nguyen, T. B., Bateman, A. P., Bones, D. L., Nizkorodov, S. A., Laskin, J., and Laskin, A.: High-resolution mass spectrometry analysis of secondary organic aerosol generated by ozonolysis of isoprene, Atmospheric Environment, 44, 1032–1042, doi:10.1016/j.atmosenv.2009.12.019, 2010.
- Nizkorodov, S. A., Laskin, J., and Laskin, A.: Molecular chemistry of organic aerosols through the application of high resolution mass spectrometry, Physical Chemistry Chemical Physics, 13, 3612–3629, doi:10.1039/c0cp02032j, 2011.
 - Noziere, B., Barnes, I., and Becker, K. H.: Product study and mechanisms of the reactions of alpha-pinene and of pinonaldehyde with OH radicals, Journal of Geophysical Research-atmospheres, 104, 23 645–23 656, doi:10.1029/1999JD900778, 1999.
- Nozière, B., Kalberer, M., Claeys, M., Allan, J., D'Anna, B., Decesari, S., Finessi, E., Glasius, M., Grgić, I., Hamilton, J. F., Hoffmann, T.,

 Iinuma, Y., Jaoui, M., Kahnt, A., Kampf, C. J., Kourtchev, I., Maenhaut, W., Marsden, N., Saarikoski, S., Schnelle-Kreis, J., Surratt, J. D.,

 Szidat, S., Szmigielski, R., and Wisthaler, A.: The Molecular Identification of Organic Compounds in the Atmosphere: State of the Art and Challenges, Chem. Rev., 115, 3919–3983, doi:10.1021/cr5003485, 2015.

- O'Boyle, N. M., Morley, C., and Hutchison, G. R.: Pybel: a Python wrapper for the OpenBabel cheminformatics toolkit, Chemistry Central Journal, 2, 5, doi:10.1186/1752-153X-2-5, 2008.
- O'Boyle, N. M., Banck, M., James, C. A., Morley, C., Vandermeersch, T., and Hutchison, G. R.: Open Babel: An open chemical toolbox, Journal of cheminformatics, 3, 1–14, doi:10.1186/1758-2946-3-33, 2011.
- Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle partitioning and secondary organic aerosol yields, Environmental Science & Technology, 30, 2580–2585, doi:10.1021/es950943+, 1996.
 - Orlan, E. S. and Boris, J. P.: Numerical Simulation of Reactive Flow, Cambridge University Press, Cambridge, U.K.; New York, NY, 2 edition edn., 2000.
- Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the atmosphere, Atmospheric Environment, 28, 185

 188, doi:10.1016/1352-2310(94)90093-0, 1994.
 - Pankow, J. F.: A consideration of the role of gas/particle partitioning in the deposition of nicotine and other tobacco smoke compounds in the respiratory tract, Chemical Research In Toxicology, 14, 1465–1481, doi:10.1021/tx0100901, 2001.
 - Pankow, J. F.: On the ability of the gas/particle partitioning constant K(p) to consider the effects of mean MW and the presence of high MW compounds, Atmospheric Environment, 45, 1213–1216, doi:10.1016/j.atmosenv.2010.11.041, 2011.
- Pankow, J. F. and Asher, W. E.: SIMPOL.1: a simple group contribution method for predicting vapor pressures and enthalpies of vaporization of multifunctional organic compounds, Atmospheric Chemistry and Physics, 8, 2773–2796, doi:10.5194/acp-8-2773-2008, 2008.
 - Pavia, D., Lampman, G., and Kriz, G.: Introduction to Spectroscopy, Brooks/Cole Pub Co., Belmont, CA, 2008.

- Pinho, P. G., Pio, C. A., Carter, W. P. L., and Jenkin, M. E.: Evaluation of alpha- and beta-pinene degradation in the detailed tro-pospheric chemistry mechanism, MCM v3.1, using environmental chamber data, Journal of Atmospheric Chemistry, 57, 171–202, doi:10.1007/s10874-007-9071-0, 2007.
- Presto, A. A. and Donahue, N. M.: Investigation of alpha-pinene plus ozone secondary organic aerosol formation at low total aerosol mass, Environmental Science & Technology, 40, 3536–3543, doi:10.1021/es052203z, 2006.
- Presto, A. A., Hartz, K. E. H., and Donahue, N. M.: Secondary organic aerosol production from terpene ozonolysis. 2. Effect of NOx concentration, Environmental Science & Technology, 39, 7046–7054, doi:10.1021/es050400s, 2005.
- Pun, B. K., Griffin, R. J., Seigneur, C., and Seinfeld, J. H.: Secondary organic aerosol 2. Thermodynamic model for gas/particle partitioning of molecular constituents, Journal of Geophysical Research-atmospheres, 107, 4333, doi:10.1029/2001JD000542, 2002.
 - Radhakrishnan, K. and Hindmarsh, A. C.: Description and use of LSODE, the Livermore solver for ordinary differential equations, Tech. Rep. UCRL-ID-113855, Lawrence Livermore National Laboratory, nASA Reference Publication 1327, 1993.
- Ranney, A. P. and Ziemann, P. J.: Microscale Spectrophotometric Methods for Quantification of Functional Groups in Oxidized Organic Aerosol, Aerosol Science and Technology, doi:10.1080/02786826.2016.1201197, 2016.
 - Reff, A., Turpin, B. J., Offenberg, J. H., Weisel, C. P., Zhang, J., Morandi, M., Stock, T., Colome, S., and Winer, A.: A functional group characterization of organic PM2.5 exposure: Results from the RIOPA study RID C-3787-2009, Atmospheric Environment, 41, 4585–4598, doi:10.1016/j.atmosenv.2007.03.054, 2007.
- Rickard, A. R., Wyche, K. P., Metzger, A., Monks, P. S., Ellis, A. M., Dommen, J., Baltensperger, U., Jenkin, M. E., and Pilling, M. J.: Gas

 phase precursors to anthropogenic secondary organic aerosol Using the Master Chemical Mechanism to probe detailed observations of
 1,3,5-trimethylbenzene photo-oxidation, Atmospheric Environment, 44, 5423–5433, doi:10.1016/j.atmosenv.2009.09.043, 2010.

- Robinson, A. L., Donahue, N. M., Shrivastava, M. K., Weitkamp, E. A., Sage, A. M., Grieshop, A. P., Lane, T. E., Pierce, J. R., and Pandis, S. N.: Rethinking organic aerosols: Semivolatile emissions and photochemical aging, Science, 315, 1259–1262, doi:10.1126/science.1133061, 2007.
- Ruggeri, G. and Takahama, S.: Technical Note: Development of chemoinformatic tools to enumerate functional groups in molecules for organic aerosol characterization, Atmospheric Chemistry and Physics Discussions, 15, 33 631–33 674, doi:10.5194/acpd-15-33631-2015, 2015.

25

- Ruppert, L., Becker, K. H., Noziere, B., and Spittler, M.: Development Of Monoterpene Oxidation Mechanisms: Results From Laboratory And Smog Chamber Studies, in: Proceedings of EUROTRAC Symposium '98, vol. 35 of WIT Transactions on Ecology and the Environment, p. 6, doi:10.2495/EURO990101, 1999.
- 10 Russell, L. M.: Aerosol organic-mass-to-organic-carbon ratio measurements, Environmental Science & Technology, 37, 2982–2987, doi:10.1021/es026123w, 2003.
 - Russell, L. M., Bahadur, R., Hawkins, L. N., Allan, J., Baumgardner, D., Quinn, P. K., and Bates, T. S.: Organic aerosol characterization by complementary measurements of chemical bonds and molecular fragments, Atmospheric Environment, 43, 6100–6105, doi:10.1016/j.atmosenv.2009.09.036, 2009.
- 15 Ruthenburg, T. C., Perlin, P. C., Liu, V., McDade, C. E., and Dillner, A. M.: Determination of organic matter and organic matter to organic carbon ratios by infrared spectroscopy with application to selected sites in the IMPROVE network, Atmospheric Environment, 86, 47–57, doi:10.1016/j.atmosenv.2013.12.034, 2014.
 - Sandu, A. and Sander, R.: Technical note: Simulating chemical systems in Fortran90 and Matlab with the Kinetic PreProcessor KPP-2.1, Atmospheric Chemistry and Physics, 6, 187–195, doi:10.5194/acp-6-187-2006, 2006.
- 20 Sato, K., Takami, A., Kato, Y., Seta, T., Fujitani, Y., Hikida, T., Shimono, A., and Imamura, T.: AMS and LC/MS analyses of SOA from the photooxidation of benzene and 1,3,5-trimethylbenzene in the presence of NO_x: effects of chemical structure on SOA aging, Atmospheric Chemistry and Physics, 12, 4667–4682, doi:10.5194/acp-12-4667-2012, 2012.
 - Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric degradation of non-aromatic volatile organic compounds, Atmospheric Chemistry and Physics, 3, 161–180, doi:10.5194/acp-3-161-2003, 2003.
 - Sax, M., Zenobi, R., Baltensperger, U., and Kalberer, M.: Time resolved infrared spectroscopic analysis of aerosol formed by photo-oxidation of 1,3,5-trimethylbenzene and alpha-pinene, Aerosol Science and Technology, 39, 822–830, doi:10.1080/02786820500257859, 2005.
 - Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, John Wiley & Sons, New York, 2nd edition edn., 2006.
- Seinfeld, J. H., Erdakos, G. B., Asher, W. E., and Pankow, J. F.: Modeling the formation of secondary organic aerosol (SOA). 2. The predicted effects of relative humidity on aerosol formation in the alpha-pinene-, beta-pinene-, sabinene-, Delta(3)-Carene-, and cyclohexene-ozone systems, Environmental Science & Technology, 35, 1806–1817, doi:10.1021/es001765+, 2001.
 - Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K. A., Seinfeld, J. H., and Pöschl, U.: Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary organic aerosol, Atmospheric Chemistry and Physics, 14, 8323–8341, doi:10.5194/acp-14-8323-2014, 2014.
 - Sindelarova, K., Granier, C., Bouarar, I., Guenther, A., Tilmes, S., Stavrakou, T., Müller, J.-F., Kuhn, U., Stefani, P., and Knorr, W.: Global data set of biogenic VOC emissions calculated by the MEGAN model over the last 30 years, Atmospheric Chemistry and Physics, 14, 9317–9341, doi:10.5194/acp-14-9317-2014, 2014.

- Subramanian, R., Khlystov, A. Y., Cabada, J. C., and Robinson, A. L.: Positive and negative artifacts in particulate organic carbon measurements with denuded and undenuded sampler configurations, Aerosol Science and Technology, 38, 27–48, doi:10.1080/02786820390229354, 2004.
- Suda, S. R., Petters, M. D., Yeh, G. K., Strollo, C., Matsunaga, A., Faulhaber, A., Ziemann, P. J., Prenni, A. J., Carrico, C. M., Sullivan,
 R. C., and Kreidenweis, S. M.: Influence of Functional Groups on Organic Aerosol Cloud Condensation Nucleus Activity, Environmental Science & Technology, doi:10.1021/es502147y, 2014.
 - Takahama, S.: APRL Substructure Search Program, doi:10.5281/zenodo.34975, https://github.com/stakahama/aprl-ssp, 2015.
 - Takahama, S. and Dillner, A. M.: Model selection for partial least squares calibration and implications for analysis of atmospheric organic aerosol samples with mid-infrared spectroscopy, Journal of Chemometrics, 29, 659–668, doi:10.1002/cem.2761, 2015.
- Takahama, S., Johnson, A., and Russell, L. M.: Quantification of Carboxylic and Carbonyl Functional Groups in Organic Aerosol Infrared Absorbance Spectra, Aerosol Science and Technology, 47, 310–325, doi:10.1080/02786826.2012.752065, 2013.
 - Tanaka, K., Kim, H.-J., Saito, K., Takahashi, H. G., Watanabe, M., Yokohata, T., Kimoto, M., Takata, K., and Yasunari, T.: How have both cultivation and warming influenced annual global isoprene and monoterpene emissions since the preindustrial era?, Atmospheric Chemistry and Physics, 12, 9703–9718, doi:10.5194/acp-12-9703-2012, 2012.
- Tolocka, M. P., Jang, M., Ginter, J. M., Cox, F. J., Kamens, R. M., and Johnston, M. V.: Formation of oligomers in secondary organic aerosol, Environmental Science & Technology, 38, 1428–1434, doi:10.1021/es035030r, 2004.
 - Topping, D., Barley, M., Bane, M. K., Higham, N., Aumont, B., Dingle, N., and McFiggans, G.: UManSysProp v1.0: an online and open-source facility for molecular property prediction and atmospheric aerosol calculations, Geoscientific Model Development, 9, 899–914, doi:10.5194/gmd-9-899-2016, 2016.
- Vayenas, D. V., Takahama, S., Davidson, C. I., and Pandis, S. N.: Simulation of the thermodynamics and removal processes in the sulfate-ammonia-nitric acid system during winter: Implications for PM2.5 control strategies, Journal of Geophysical Research-atmospheres, 110, doi:10.1029/2004JD005038, 2005.
 - Wang, Y., Kim, H., and Paulson, S. E.: Hydrogen peroxide generation from α and β -pinene and toluene secondary organic aerosols, Atmospheric Environment, 45, 3149 3156, doi:10.1016/j.atmosenv.2011.02.060, 2011.
- Wyche, K. P., Monks, P. S., Ellis, A. M., Cordell, R. L., Parker, A. E., Whyte, C., Metzger, A., Dommen, J., Duplissy, J., Prevot, A. S. H., Baltensperger, U., Rickard, A. R., and Wulfert, F.: Gas phase precursors to anthropogenic secondary organic aerosol: detailed observations of 1,3,5-trimethylbenzene photooxidation, Atmospheric Chemistry and Physics, 9, 635–665, 2009.
 - Yanenko, N. N.: The Method of Fractional Steps: The Solution of Problems of Mathematical Physics in Several Variables, Springer, 1 edition edn., 1971.
- 30 Yeh, G. K. and Ziemann, P. J.: Gas-Wall Partitioning of Oxygenated Organic Compounds: Measurements, Structure–Activity Relationships, and Correlation with Gas Chromatographic Retention Factor, Aerosol Science and Technology, 49, 727–738, doi:10.1080/02786826.2015.1068427, 2015.
 - Yu, J. Z., Cocker, D. R., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: Gas-phase ozone oxidation of monoterpenes: Gaseous and particulate products, Journal of Atmospheric Chemistry, 34, 207–258, doi:10.1023/A:1006254930583, 1999.
- Zeng, G., Holladay, S., Langlois, D., Zhang, Y., and Liu, Y.: Kinetics of Heterogeneous Reaction of Ozone with Linoleic Acid and its Dependence on Temperature, Physical State, RH, and Ozone Concentration, The Journal of Physical Chemistry A, 117, 1963–1974, doi:10.1021/jp308304n, 2013.

- Zhang, X. and McMurry, P.: Theoretical analysis of evaporative losses from impactor and filter deposits, Atmospheric Environment (1967), 21, 1779 1789, doi:10.1016/0004-6981(87)90118-1, 1987.
- Zhang, X. and Seinfeld, J. H.: A functional group oxidation model (FGOM) for SOA formation and aging, Atmospheric Chemistry and Physics, 13, 5907–5926, doi:10.5194/acp-13-5907-2013, 2013.
- 5 Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of secondary organic aerosol, Proceedings of the National Academy of Sciences of the United States of America, 111, 5802–5807, doi:10.1073/pnas.1404727111, 2014a.
 - Zhang, X., Schwantes, R. H., Coggon, M. M., Loza, C. L., Schilling, K. A., Flagan, R. C., and Seinfeld, J. H.: Role of ozone in SOA formation from alkane photooxidation, Atmospheric Chemistry and Physics, 14, 1733–1753, doi:10.5194/acp-14-1733-2014, 2014b.
- Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and Seinfeld, J. H.: Formation and evolution of molecular products in α-pinene secondary organic aerosol, Proceedings of the National Academy of Sciences, 112, 14168–14173, doi:10.1073/pnas.1517742112, 2015.
 - Ziemann, P. J. and Atkinson, R.: Kinetics, products, and mechanisms of secondary organic aerosol formation, Chemical Society Reviews, 41, 6582–6605, doi:10.1039/c2cs35122f, 2012.
- Zuend, A., Marcolli, C., Booth, A. M., Lienhard, D. M., Soonsin, V., Krieger, U. K., Topping, D. O., McFiggans, G., Peter, T., and Seinfeld, J. H.: New and extended parameterization of the thermodynamic model AIOMFAC: calculation of activity coefficients for organic-inorganic mixtures containing carboxyl, hydroxyl, carbonyl, ether, ester, alkenyl, alkyl, and aromatic functional groups, Atmospheric Chemistry and Physics, 11, 9155–9206, doi:10.5194/acp-11-9155-2011, 2011.

Tables

Table 1. Summary of the experimental conditions studied in this work. For simplification, an ID has been given to each system.

ID	Publication	Precursor	Measurement conditions	
APIN-lNO _x	Sax et al., 2005		low NO _x : 240 ppb	
		α -pinene: 300 ppb	RH: 61%	
			seed: none	
			radical initiator: propene, 300 ppb	
APIN-hNO _x	Chhabra et al., 2011		high NO _x : 847 ppb	
		α -pinene: 47 ppb reacted	RH: 5%	
			seed: ammonium sulfate, 27 $\mu \mathrm{g} \ \mathrm{m}^{-3}$	
			radical initiator: CH ₃ ONO, 200-400 ppb	
APIN-nNO _x	Chhabra et al., 2011	α -pinene: 46 ppb reacted	no NO _x	
			RH: 4%	
			seed: ammonium sulfate, 24 $\mu \mathrm{g} \ \mathrm{m}^{-3}$	
			radical initiator: H ₂ O ₂ ,	
TMB-lNO _x	Sax et al., 2005	1,3,5-trimethyl benzene: 1312 ppb	low NO _x : 320 ppb	
			RH: 60%	
			seed: none	
			radical initiator: propene, 300 ppb	

Table 2. Illustration of several polyfunctional molecules discussed in Section 3.

MCM name	Molecular weight	$\log C^0$ (298 K)	Structure
Н3С25ССО2Н	174.1513	1.16	H₃C OH OH OH
С719ООН	176.1672	0.97	OH CH ₃ OH
С97ООН	188.2209	2.33	H ₃ C CH ₃ OH
С98ООН	204.2203	1.45	OH CH ₃ OCH ₃
С106ООН	216.2310	1.92	HO O CH ₃ CCH ₃
тм135ВРООН	202.2045	2.78	H ₃ C O CH ₃ H ₃ C
NMXYFUOOH	207.1382	3.40	O,N-O CH ₃ OH H ₃ C-O
C813NO3	235.1913	-0.38	ONO CH3 OHO OH
C811PAN	247.2020	2.17 26	H ₃ C CH ₃ O O N O

Table 3. Comparison of pure component vapor pressures (atm) estimated (at 298 K) for the most abundant ONO₂ compounds in the aerosol phase for the APIN-hNO_x simulation. Calculations were accessed using the UManSysProp tool (Toppi al., 2016).

Compound	SIMPOL.1 ¹	EVAPORATION ²	Nannoolal ³	Myrdal & Yalkowsky ⁴
C813NO3	4.33×10^{-11}	4.04×10^{-11}	1.70×10^{-11}	7.05×10^{-9}
C98NO3	6.12×10^{-9}	2.07×10^{-8}	8.15×10^{-9}	1.01×10^{-7}
C719NO3	2.33×10^{-9}	3.34×10^{-10}	5.96×10^{-9}	3.45×10^{-7}
APINANO3	1.55×10^{-7}	2.38×10^{-6}	9.47×10^{-7}	5.53×10^{-6}
APINBNO3	1.55×10^{-7}	8.19×10^{-6}	1.45×10^{-6}	7.39×10^{-6}
TM135BP	7.22×10^{-8}	5.46×10^{-9}	8.42×10^{-11}	6.24×10^{-9}
NMXYFUOOH	2.96×10^{-7}	4.56×10^{-8}	3.41×10^{-9}	5.09×10^{-8}

¹Pankow and Asher (2008)

²Compernolle et al. (2011)

³Nannoolal et al. (2008)

⁴Myrdal and Yalkowsky (1997)

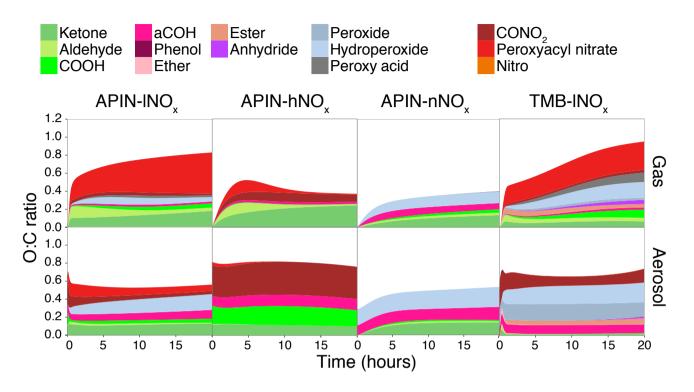


Figure 1. Time series of the relative molar contribution of different FGs to the O:C in the gas phase (top panel) and aerosol phase (bottom panel) simulated in this work for APIN-INO $_x$, APIN-hNO $_x$, APIN-nNO $_x$, and TMB-INO $_x$. The contribution of each FG to the O:C ratio accounts for the number of oxygen atoms per FG.

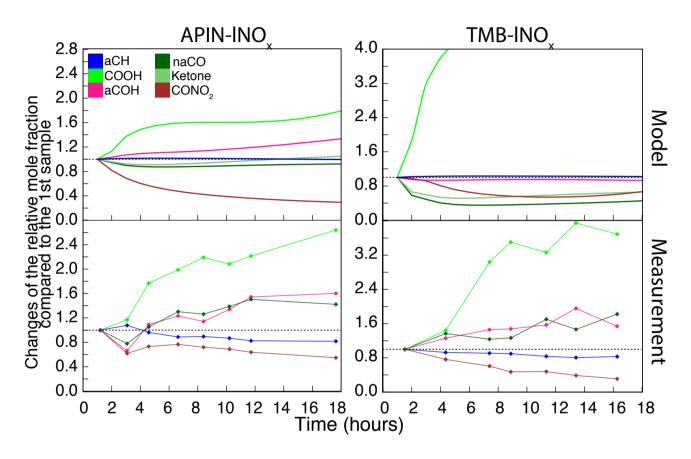


Figure 2. Comparison in the changes of the relative mole fraction compared to the first sample for COOH, CO, aCH and CONO₂ of the aerosol phase measured by Sax et al. (2005) and modeled in this work for APIN-lNO_x and TMB-lNO_x. For a chosen FG, the changes of the relative mole fraction compared to the first sample is calculated as the ratio between the relative mole fraction at the chosen time and the relative mole fraction at 1 hour. naCO includes ketone and aldehyde FGs, but the change in relative ketone FG abundance is also shown separately for illustration. The contribution of ketone and aldehyde to CO have been reported separately in the model results. The x-axis refers to the hours after the lights were turned on in the chamber for the top panel (Measurement), and the time after the start of the simulation in the bottom panel (Model). The dashed line corresponds to y=1 and has been added for visual reference.

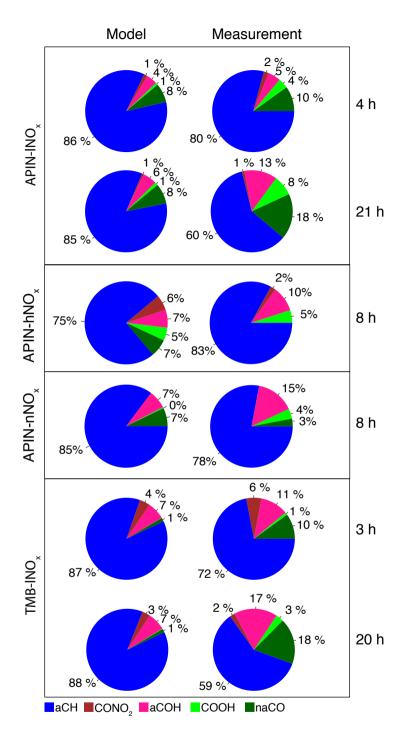


Figure 3. Pie charts illustrating the time-integrated relative aerosol mole fraction of aCH, CO, COOH, CONO₂, and aCOH in model simulations and experiments. The mole fractions reported in simulations are summed with respect to the subset of FGs that are reported by measurement to facilitate direct comparison. The time reported refers to the hours after the lights were turned on in the chamber (Measurements), and the time after the start of the simulation (Model). In the pie charts reporting the measurement conducted by Chhabra et al. (2011) (APIN-hNO_x and APIN-nNO_x) the CNH₂ fraction has been omitted 30 order to obtain a direct comparison between model and experiments. The sum of percentages combines to $100\pm1\%$, as individual values were rounded to the nearest whole number for labeling.

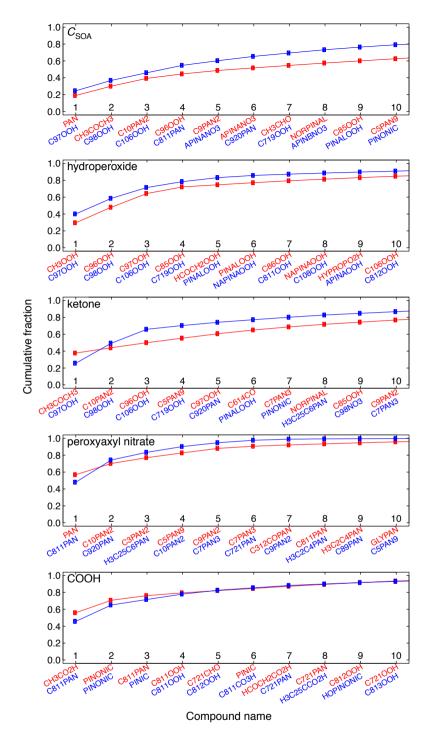


Figure 4. Cumulative contribution (as a fraction of total) of each compound to the overall C_{SOA} mass and abundance of different FG fragments for the APIN-lNO_x simulation. Compounds are arranged in order of decreasing contribution in each phase (i.e., first molecule contributes the greatest amount). Contributions to the aerosol phase are shown in blue and the gas phase in red.

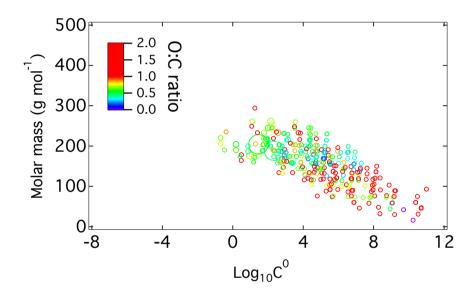


Figure 5. Molar mass vs the logarithm of the pure components saturation concentration for the compounds included in the APIN and propene MCMv3.2 degradation scheme. The size of the circles is proportional to the compound mass in the aerosol phase found in APIN- lNO_x simulation.

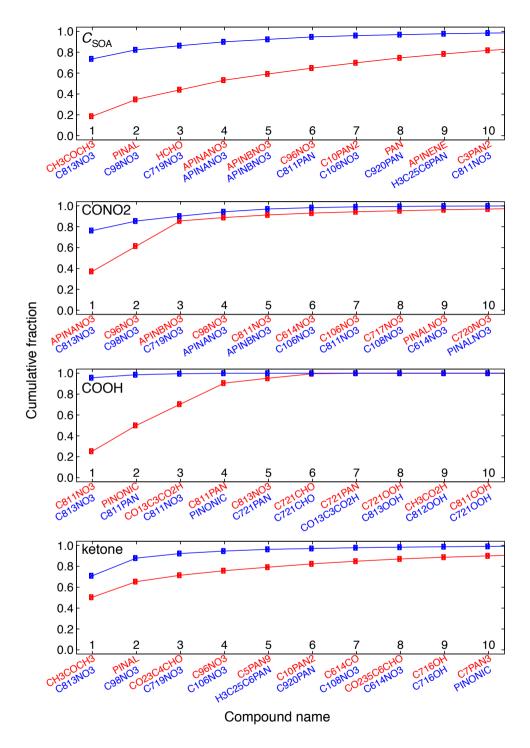


Figure 6. Cumulative contribution (as a fraction of total) of each compound to the overall C_{SOA} mass and abundance of different FG fragments for the APIN-hNO_x simulation. Compounds are arranged in order of decreasing contribution (i.e., the first compound contributes most) for each phase. Contributions to the aerosol phase are shown in blue and the gas phase in red.

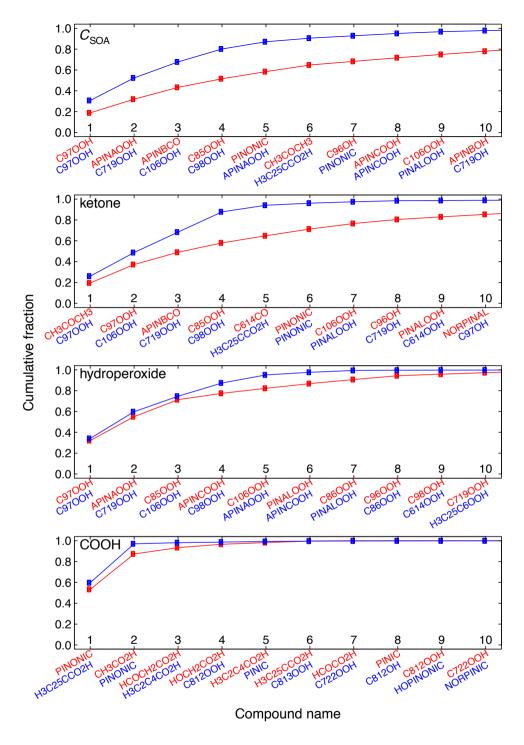


Figure 7. Cumulative contribution (as a fraction of total) of each compound to the overall C_{SOA} mass and abundance of different FGs for the APIN-nNO_x simulation. Componds are arranged in order of decreasing contribution (i.e., the first compound contributes most) for each phase. Contributions to the aerosol phase are shown in blue and the gas phase in red.

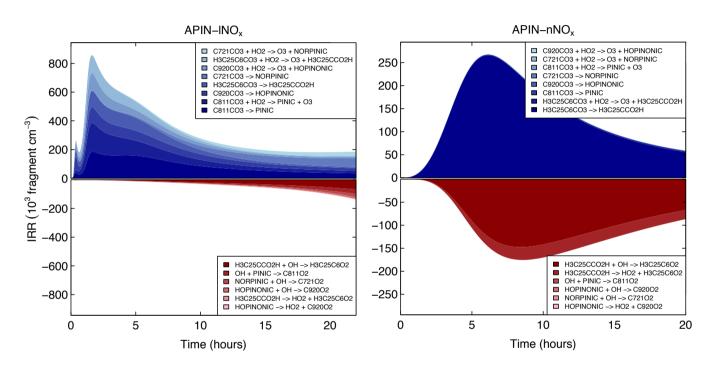


Figure 8. Integrated Reaction Rates for the COOH group (denoted in units of fragments per molecule).

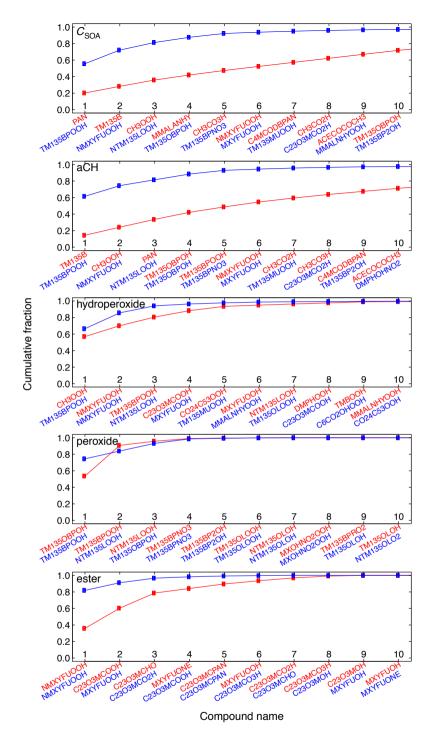


Figure 9. Cumulative contribution (as a fraction of total) of each compound to the overall C_{SOA} mass and abundance of different FGs for the TMB-lNO_x simulation. Componds are arranged in order of decreasing contribution (i.e., the first compound contributes most) for each phase. Contributions to the aerosol phase are shown in blue and the gas phase in red.

Supplemental material for "Model-measurement comparison of functional group abundance in α -pinene and 1,3,5-trimethylbenzene secondary organic aerosol formation"

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This document contains Appendices A and B as well as supplemental figures S1–S14.

Appendix A: Dynamic absorptive partitioning

In this section, we summarize the dynamic absorptive partitioning batch reactor model (inspired by the continuous flow reactor model of Chen et al. (2011)) and also our use of pure component and saturation concentrations. Symbols and their descriptions are listed in Table A1.

We first revisit the relationship among equilibrium vapor pressure p of a substance i in a mixture, its pure component vapor pressure (possibly over sub-cooled liquid) p^0 , and its activity a of solution. The vapor and aerosol solution phase chemical potentials can be written as a sum of their standard chemical potentials μ^0 and μ^* and their ideal (and non-ideal, in the case of liquid) mixing contributions to the partial molar Gibbs free energy (Denbigh, 1981; Seinfeld and Pandis, 2006):

$$\mu_i^{\text{(vap.)}} = \mu_i^0 + RT \ln p_i \tag{A1}$$

$$\mu_i^{\text{(soln.)}} = \mu_i^* + RT \ln a_i . \tag{A2}$$

At equilibrium, the chemical potential of substance i in the two phases are equal, $\mu_i^{\text{(vap.)}} = \mu_i^{\text{(soln.)}}$; the equilibrium constant K is given by the relation between equations A1 and A2:

15
$$K_i = \exp\left(\frac{\mu_i^* - \mu_i^0}{RT}\right) = \frac{p_i}{a_i}$$
 (A3)

Using a pure component reference (as opposed to infinite dilution) for all species, $K_i = p_{L,i}^0$. With activity of substance i defined as $a_i = \zeta_i x_i$, the equality in equation A3 is commonly written as:

$$p_i = \zeta_i x_i p_{L,i}^0 . (A4)$$

The pure component and equilibrium vapor pressures are equivalent to their mass concentrations C^0 and $C^{(g)}$, respectively at a given temperature by the ideal gas law (Chen et al., 2011; Donahue et al.,

Table A1. Summary of variables and notation used for model description.

Symbol	Description
$\overline{}$	compound index
n	number of moles
x	mole fraction in solution
μ	chemical potential ($kJ \text{ mol}^{-1}$)
μ^0	standard chemical potential (gas-phase reference at 1 atm, $\mathrm{kJ}\mathrm{mol}^{-1}$)
μ^*	standard chemical potential (liquid-phase reference, $\mathrm{kJ}\mathrm{mol}^{-1})$
K	equilibrium constant
a	solution activity
ζ	activity coefficient
p	equilibrium vapor pressure (atm)
p_L^0	pure component vapor pressure (atm)
$C^{(g)}$	gas-phase concentration ($\mu g m^{-3}$)
$C^{(p)}$	particle-phase concentration ($\mu g \mathrm{m}^{-3}$)
C^0	pure component concentration ($\mu g m^{-3}$)
C^*	saturation concentration ($\mu g m^{-3}$)
C_{OA}	organic aerosol concentration ($\mu g m^{-3}$)
MW	molecular weight $(g \text{ mol}^{-1})$
\overline{MW}	mean molecular weight $(g \text{ mol}^{-1})$
R	gas constant $(m^3 atm K^{-1} mol^{-1})$
T	temperature (K)
K_p	partitioning coefficient ($m^3 \mu g^{-1}$)
D_p	particle diameter (nm)
D_{seed}	seed diameter (nm)
D	diffusivity (m^2s^{-1})
N_p	number concentration (m ⁻³)
$ ho_p$	SOA density $(g cm^{-3})$
Kn	Knudsen number
λ	mean free path (m)

2012):

$$C_i^0 = \frac{MW_i}{RT} p_{L,i}^0 \quad \text{and} \quad C_i^{(g)} = \frac{MW_i}{RT} p_i \ . \tag{A5} \label{eq:A5}$$

The saturation concentration C^* (Donahue et al., 2006) is also a widely used metric for characterizing the volatility of a mixture; it is the reciprocal of the venerable G/P partition coefficient K_p

25 (Pankow, 1994, 2011):

$$C_i^* = \frac{C_i^{(g)}}{C_i^{(p)}} C_{\text{OA}} = \left(\frac{C_i^{(p)}}{C_{\text{OA}}}\right)^{-1} C_i^{(g)} = K_{p,i}^{-1} . \tag{A6}$$

The mass fraction of i in the particle phase (with respect to the total organic aerosol mass) can be expressed according to its molar abundance n, mole fraction x, and molecular weight \overline{MW} defined for the set of compounds $\mathcal M$ in the mixture:

30
$$\frac{C_i^{(p)}}{C_{\text{OA}}} = \frac{MW_i n_i}{\sum_{i \in \mathcal{M}} MW_i n_i} = \frac{MW_i}{\overline{MW}} \cdot \frac{n_i}{\sum_{i \in \mathcal{M}} n_i} = \frac{MW_i}{\overline{MW}} x_i . \tag{A7}$$

By substitution of equation A7 into A4–A6, C_i^* can be related to C_i^0 :

$$C_i^* = \frac{\overline{MW}}{RT} \frac{p_i}{x_i} = \frac{\overline{MW}}{RT} \zeta_i p_{L,i}^0 = \frac{\overline{MW}}{MW_i} \zeta_i C_i^0$$
(A8)

For a pure component solution, $C_i^* = C_i^0$. At equilibrium with a solution mixture, the gas phase concentration is related to saturation concentrations C_i^0 and C_i^* through equations A4–A6:

35
$$C_i^{(g)} = a_i C_i^0 = \frac{C_i^{(p)}}{C_{OA}} C_i^*$$

 C^* is a useful construct to represent volatility of aggregated mixtures when its composition is not well-defined (Donahue et al., 2006; Grieshop et al., 2009; Chen et al., 2011), and we use this for specifying initial concentrations of a generic organic aerosol mixture (Appendix B). For subsequent G/P partitioning calculations, we follow the approach of Donahue et al. (2012) and use C^0 according to molecular specificity provided by the MCMv3.2 mechanism.

Considering the dynamics of condensation and evaporation, mass transfer of individual compounds from the bulk vapor phase (at concentration $C_{i,\infty}^{(g)}$) to a monodisperse particle population of size D_p is driven by the concentration gradient with respect to the equilibrium vapor concentration at the aerosol surface, and the surface characteristics for exchange to take place (Fuchs and Sutugin, 1971; Seinfeld and Pandis, 2006; Chen et al., 2011):

$$\frac{dC_{i,\infty}^{(g)}}{dt} = -\frac{dC_{i}^{(p)}}{dt} = -2\pi N_p D_p D_i f(\alpha_i, Kn) (C_{i,\infty}^{(g)} - a_i C_i^0)
f(\alpha_i, Kn) = \frac{0.75\alpha_i (1 + Kn)}{Kn^2 + Kn + 0.283Kn\alpha_i + 0.75\alpha_i}$$
(A9)

We assume activity coefficients ζ_i and mass accommodations α_i of unity, and a diffusivity of $D_i = 5 \times 10^{-6} \ \mathrm{m^2 \, s^{-1}}$ for all molecules (Chen et al., 2011). One notable difference with Chen et al. (2011) is that we run this model as a batch reactor rather than a continuously-stirred flow tank reactor, in accordance with the operation of experimental chambers of the reference studies used in this work.

Given our assumption of monodisperse particles and fixed particle number N_p (no particle losses; negligible coagulation for the duration of our simulations), we estimate the size D_p of particles

(which includes the seed, if applicable, and condensed organic matter) at each timestep to satisfy the following condition, assuming SOA density of $\rho_p = 1.5 \text{ g cm}^{-3}$ (Kostenidou et al., 2007):

$$C_{\rm OA} = N_p \rho_p \frac{\pi}{6} \left(D_p^3 - D_{\rm seed}^3 \right)$$

60

Values for these parameters are discussed in Appendix B.

Appendix B: Initial and fixed conditions

In addition to precursor, $\mathrm{NO_x}$, and oxidant concentrations specified in Table 1, we initialize gasphase concentrations to small, non-zero values $(C_i^{(g)})$ set equivalent to 10^{-5} ppb for each compound) to prevent numerical singularities. The initial concentration of the aerosol phase is important also as it determines the activity (i.e., mole fraction) in equation A9. Assuming values at either extreme of the domain can lead to undesired behavior. In the perfect sink assumption (all $x_i = 0$), excess condensation of the most abundant gas-phase product (without regard to compound-specific saturation concentration) can occur in a single time step. In the pure component or external mixture assumption (all $x_i = 1$), condensation during initial stages can be extremely slow for simulations under low precursor concentrations, and erratic changes in equilibrium concentrations can occur when transitioning from $x_i = 1$ to $x_i \ll 1$ immediately following initial condensation.

The use of an organic seed aerosol ($C_{\text{OA,init}}$) to initiate G/P partitioning alleviates such problems and accounts for the effects of homogeneous and heterogeneous nucleation (Flagan and Seinfeld, 1988; Seinfeld and Pandis, 2006; Fan et al., 2013, e.g.,) of organic aerosol not included in our model. Pankow (2001) used $C_{\text{OA,init}}$ and composition (x_i s) from the final C_{OA} values determined from experiments. As this can introduce additional mass into the system, we introduce a different approach. Initial aerosol concentrations are specified to be in equilibrium with the specified gasphase concentrations for a given $C_{\text{OA,init}}$ (according to equations A6 and A8 with $\zeta_i = 1$):

$$C_i^{(p)} = \left(\frac{C_i^{(g)}}{C_i^*}\right) C_{\text{OA,init}} = \left(\frac{C_i^{(g)}}{C_i^0} \cdot \frac{MW_i}{\overline{MW}}\right) C_{\text{OA,init}} . \tag{B1}$$

This leads to the introduction of negligible mass of MCM compounds into the system, but provides a means for determining non-zero initial values of x_i s that scale with the vapor pressures of condensible products. For typical values of $C_{\text{OA,init}}$, $\sum_{i \in \mathcal{M}'} C_i^{(p)} \ll C_{\text{OA,init}}$ where \mathcal{M}' represents all compounds in the MCM mechanism; we assume the existence of an inactive medium (C_{solvent}) which comprises this difference:

$$C_{\text{solvent}} = C_{\text{OA,init}} - \sum_{i \in \mathcal{M}'} C_i^{(p)}$$
.

 C_{solvent} does not participate in G/P partitioning or reactions, and is not reported with the SOA formed. The initial and subsequent mole fractions are calculated in the presence of this virtual medium:

85
$$x_i = \frac{C_i^{(p)}/MW_i}{\sum_{i \in \mathcal{M}'} C_i^{(p)}/MW_i + C_{\text{solvent}}/\overline{MW_{\text{solvent}}}}$$
 (B2)

 \mathcal{M}' , in combination with this inactive medium, comprises the complete set of molecule types \mathcal{M} in the system (Appendix A).

A value of $C_{\text{OA,init}} = 1 \, \mu \mathrm{g \, m^{-3}}$ is assumed in all of our simulations, which leads to initial values of $\sum_i C_i^{(p)} < 5 \times 10^{-4} \, \mu \mathrm{g \, m^{-3}}$. Therefore, $C_{\text{solvent}} \sim C_{\text{OA,init}}$ and the organic aerosol phase affecting G/P partitioning is effectively $C_{\text{OA}} = C_{\text{solvent}} + C_{\text{SOA}} \approx C_{\text{OA,init}} + C_{\text{SOA}}$. For use in both equations B1 and B2 we assume a mean molecular weight of $\overline{MW} = \overline{MW_{\text{solvent}}} = 200 \, \mathrm{g \, mol^{-1}}$, which is in the range of products estimated or identified in the condensed phase (e.g., Odum et al., 1996; Nguyen et al., 2010; Shiraiwa et al., 2014). While C_{SOA} is several orders of magnitude greater than $C_{\text{OA,init}} = 1 \, \mu \mathrm{g \, m^{-3}}$ for APIN-lNO_x and TMB-lNO_x simulations (and therefore dominates x_i s), C_{SOA} produced for APIN-hNO_x and APIN-nNO_x is within the same order of magnitude for conceivable values of $C_{\text{OA,init}}$ near unity. However, in Figures S3 and S4 we illustrate that relative abundances of FGs analyzed reported in this work are relatively insensitive to the value of $C_{\text{OA,init}}$ and C_{SOA} for these latter simulations.

For experiments with ammonium sulfate seed (APIN-hNO $_x$ and APIN-nNO $_x$), we assume $D_{\rm seed}=200$ nm and $N_p=3.5\times10^3~{\rm cm^{-3}}$, which corresponds to approximately 26 $\mu{\rm g\,m^{-3}}$ of ammonium sulfate in accordance with experiments (Table 1). With $C_{\rm OA,init}=1\mu{\rm g\,m^{-3}}$ and $\rho_p=1.5~{\rm g\,cm^{-3}}$, initial D_p becomes 203 nm. For APIN-lNO $_x$ and TMB-lNO $_x$ where no seed is used, we assume $N_p=10^4~{\rm cm^{-3}}$ and initial $D_p=43$ nm is calculated according to $C_{\rm OA,init}$, which corresponds to a system which condensational growth and scavenging of the smallest clusters has occured. The Kelvin effect is neglected but results are insensitive to this omission since the condensation and particle growth is rapid in these two systems. D_p increases up to 350 nm over the course of the APIN-lNO $_x$ and TMB-lNO $_x$ simulations for the given $C_{\rm OA}$ loadings, and up to 220 nm (including seed) for APIN-hNO $_x$ and APIN-nNO $_x$ on account of the lower mass loading. Inorganic reactions (such as the production of nitric acid) in simulations involving NO $_x$ are included to maintain radical balance, but are not treated in G/P partitioning. Rate of particle growth may therefore be underestimated when condensation of such species (e.g., nitric acid) may be important compared to the growth due to the organic phase.

Sensitivity of the rate of mass transfer (equation A9) to N_p , D_p , $\overline{MW}_{\text{solvent}}$ described above can be interpreted through the sensitivity to $C_{\text{OA,init}}$ as demonstrated for the APIN-hNO_x and APIN-nNO_x simulations. However, the chemistry of the gas phase and formation of condensible products is not affected significantly as the mass removed to the aerosol fraction (where they are shielded from further chemical reaction in our model) is generally small (Figure S12). Therefore, the relative composition of the gas and aerosol phases analyzed in this work are not changed significantly by these parameters in our simulations.

120 Supplemental figures

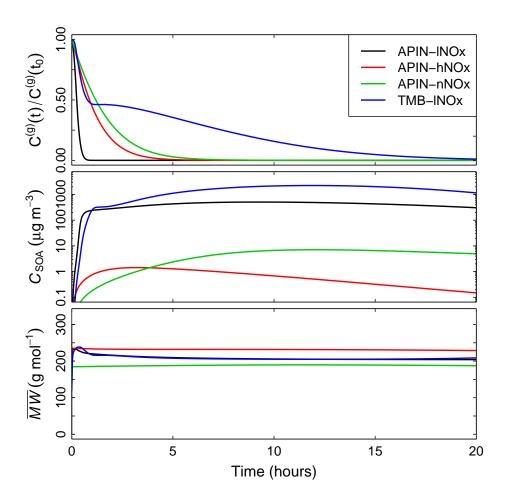


Figure S1. Time series of precursor concentrations ratioed to initial value (top), C_{SOA} (= $C_{OA} - C_{OA,init}$) concentrations (middle), and mean molecular weight of compounds weighted by abundance in the aerosol phase (bottom).

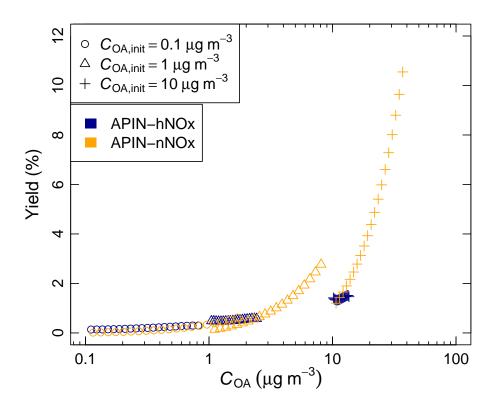


Figure S2. Aerosol yield curves for various values of $C_{\text{OA,init}}$ in the APIN-hNO_x and APIN-nNO_x simulations. $C_{\text{OA}} (= C_{\text{OA,init}} + C_{\text{SOA}})$, which includes the SOA formed and $C_{\text{OA,init}}$ to initiate G/P partitioning.

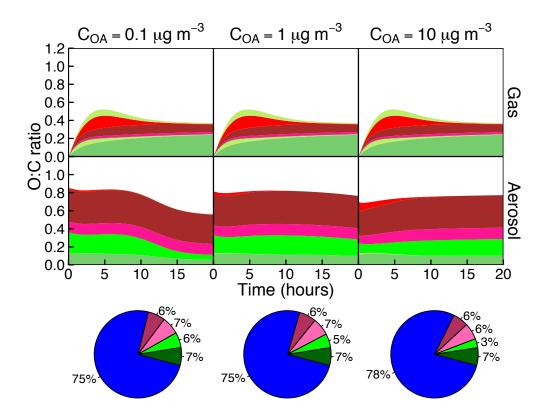


Figure S3. Evolution of O:C ratio for various values of $C_{OA,init}$ in the APIN-hNO_x simulation. Color schemes are same as for Figures 1 and 3. Pie charts illustrate time-integrated composition for the first 8 hours.

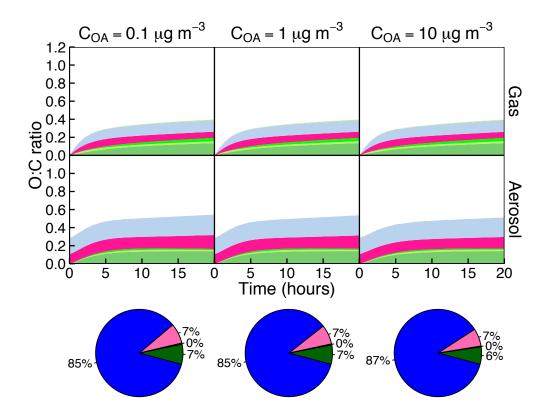


Figure S4. Evolution of O:C ratio for various values of $C_{OA,init}$ in the APIN-nNO_x simulation. Color schemes are same as for Figures 1 and 3. Pie charts illustrate time-integrated composition for the first 8 hours.

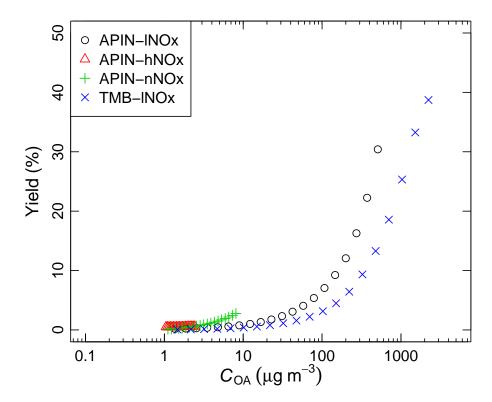


Figure S5. Aerosol yields ($C_{\rm OA,init}$ =1 $\mu \rm g \ m^{-3}$) as a function of $C_{\rm OA}$ (= $C_{\rm OA,init}$ + $C_{\rm SOA}$), which includes the SOA formed and $C_{\rm OA,init}$ to initiate G/P partitioning.

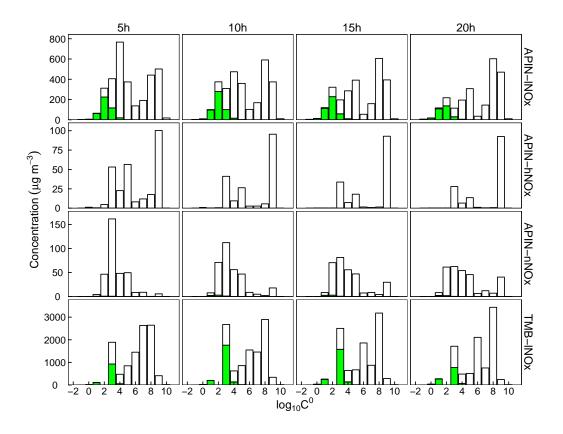


Figure S6. Volatility distribution of gas (white) and aerosol (green) phase products at different time intervals of the photooxidation simulation ($C_{OA,init}=1~\mu g~m^{-3}$). C^0 is in units of $\mu g~m^{-3}$. Different simulations are shown in different rows with varying scales along y-axis.

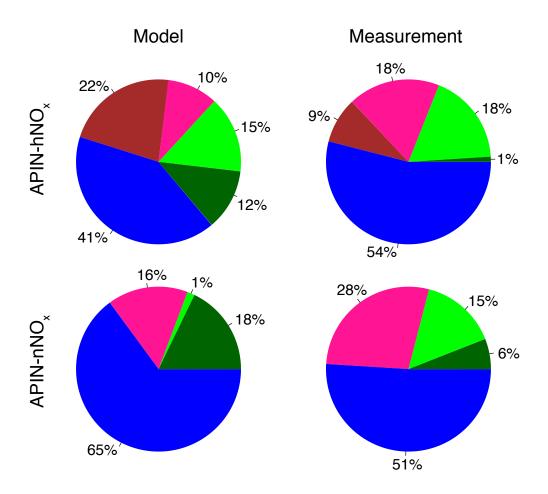


Figure S7. Relative mass abundance of measured FGs estimated by converting molar abundance of FG fragments (Figure 3) to their mass equivalents using coefficients assumed by Chhabra et al. (2011).

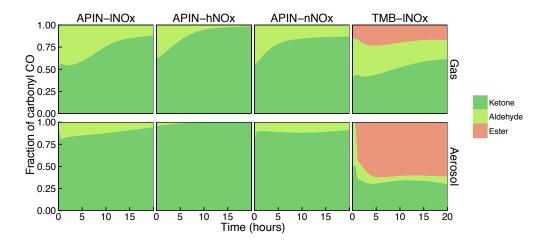


Figure S8. Fraction of carbonyl CO in each phase.

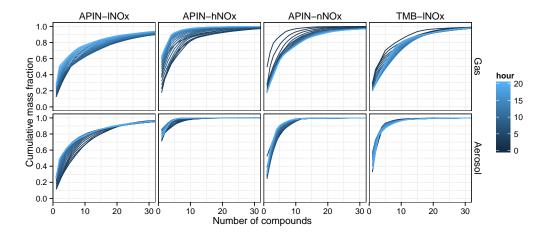


Figure S9. Cumulative mass fraction of compounds in the gas (top row) and aerosol (bottom row) phases for each simulations (shown along columns). For each phase, compounds are arranged in descending order according to their contribution (largest first).

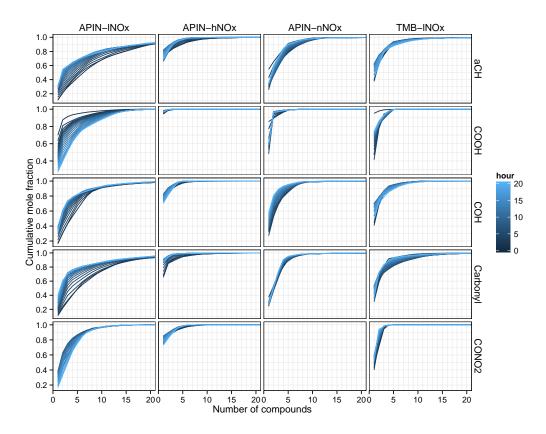


Figure S10. Cumulative fraction of compounds contributing to each FG (shown along rows) in the aerosol phase. For each FG, compounds are arranged in descending order according to their contribution (largest first). The lower bounds of y-axes on different rows are greater than zero, and are selected such that gradations in lines can be better differentiated.

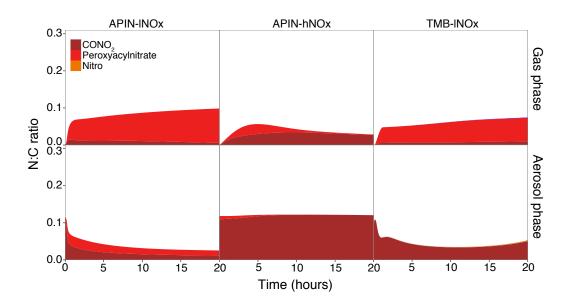


Figure S11. FGs contribution to the total N:C ratio for each simulation in which nitrogen is present. The only nitrogen containing FGs present in the system are CONO₂, peroxyacyl nitrate and nitro (amines and amides are not included).

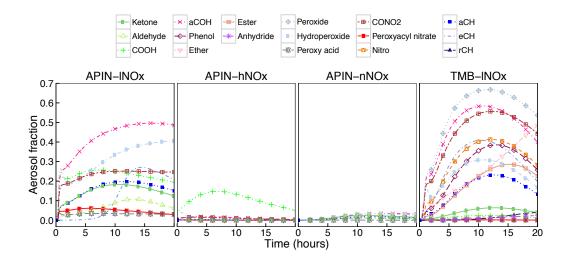


Figure S12. Evolution of aerosol fraction for each FG.

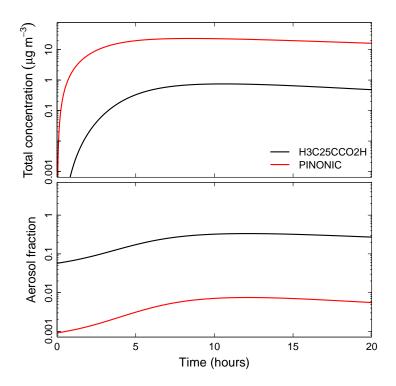


Figure S13. Total (gas+aerosol) concentrations (top) and aerosol fractions (bottom) of the two dominant acids in the APIN- nNO_x system.

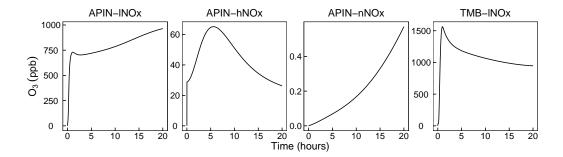


Figure S14. Time series of ozone concentrations in each of the four simulations studied.

References

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- Chen, Q., Liu, Y., Donahue, N. M., Shilling, J. E., and Martin, S. T.: Particle-Phase Chemistry of Secondary Organic Material: Modeled Compared to Measured O:C and H:C Elemental Ratios Provide Constraints, Environmental Science & Technology, 45, 4763–4770, doi:10.1021/es104398s, 2011.
- 125 Chhabra, P. S., Ng, N. L., Canagaratna, M. R., Corrigan, A. L., Russell, L. M., Worsnop, D. R., Flagan, R. C., and Seinfeld, J. H.: Elemental composition and oxidation of chamber organic aerosol, Atmospheric Chemistry and Physics, 11, 8827–8845, doi:10.5194/acp-11-8827-2011, 2011.
 - Denbigh, K.: The Principles of Chemical Equilibrium, Cambridge University Press, Cambridge, UK, 1981.
- Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning, dilution, and chemical aging of semivolatile organics, Environmental Science & Technology, 40, 2635–2643, doi:10.1021/es052297c, 2006.
 - Donahue, N. M., Kroll, J. H., Pandis, S. N., and Robinson, A. L.: A two-dimensional volatility basis set Part 2: Diagnostics of organic-aerosol evolution, Atmospheric Chemistry and Physics, 12, 615–634, doi:10.5194/acp-12-615-2012, 2012.
- 135 Fan, Y., Qin, F., Luo, X., Lin, L., Gui, H., and Liu, J.: Heterogeneous condensation on insoluble spherical particles: Modeling and parametric study, Chemical Engineering Science, 102, 387 396, doi:10.1016/j.ces.2013.08.040, 2013.
 - Flagan, R. C. and Seinfeld, J. H.: Fundamentals of air pollution engineering, Prentice-Hall, Inc., Englewood Cliffs, New Jersey, 1988.
- 140 Fuchs, N. A. and Sutugin, A. G.: High-Dispersed {Aerosols}, in: Topics in Current Aerosol Research, edited by Brock, G. M. and Hidy, J. R., International Reviews in Aerosol Physics and Chemistry, pp. 1–3, Pergamon, doi:10.1016/B978-0-08-016674-2.50006-6, 1971.
 - Grieshop, A. P., Miracolo, M. A., Donahue, N. M., and Robinson, A. L.: Constraining the Volatility Distribution and Gas-Particle Partitioning of Combustion Aerosols Using Isothermal Dilution and Thermodenuder Measurements, Environmental Science & Technology, 43, 4750–4756, doi:10.1021/es8032378, 2009.
 - Kostenidou, E., Pathak, R. K., and Pandis, S. N.: An algorithm for the calculation of secondary organic aerosol density combining AMS and SMPS data, Aerosol Science and Technology, 41, 1002–1010, doi:10.1080/02786820701666270, 2007.
- Nguyen, T. B., Bateman, A. P., Bones, D. L., Nizkorodov, S. A., Laskin, J., and Laskin, A.: High-resolution mass spectrometry analysis of secondary organic aerosol generated by ozonolysis of isoprene, Atmospheric Environment, 44, 1032–1042, doi:10.1016/j.atmosenv.2009.12.019, 2010.
 - Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.: Gas/particle partitioning and secondary organic aerosol yields, Environmental Science & Technology, 30, 2580–2585, doi:10.1021/es950943+, 1996.
- Pankow, J. F.: An absorption model of gas/particle partitioning of organic compounds in the atmosphere, Atmospheric Environment, 28, 185 188, doi:10.1016/1352-2310(94)90093-0, 1994.
 - Pankow, J. F.: A consideration of the role of gas/particle partitioning in the deposition of nicotine and other tobacco smoke compounds in the respiratory tract, Chemical Research In Toxicology, 14, 1465–1481, doi:10.1021/tx0100901, 2001.

- 160 Pankow, J. F.: On the ability of the gas/particle partitioning constant K(p) to consider the effects of mean MW and the presence of high MW compounds, Atmospheric Environment, 45, 1213–1216, doi:10.1016/j.atmosenv.2010.11.041, 2011.
 - Seinfeld, J. H. and Pandis, S. N.: Atmospheric Chemistry and Physics: From Air Pollution to Climate Change, John Wiley & Sons, New York, 2nd edition edn., 2006.
- Shiraiwa, M., Berkemeier, T., Schilling-Fahnestock, K. A., Seinfeld, J. H., and Pöschl, U.: Molecular corridors and kinetic regimes in the multiphase chemical evolution of secondary organic aerosol, Atmospheric Chemistry and Physics, 14, 8323–8341, doi:10.5194/acp-14-8323-2014, 2014.