

Interactive comment on “New Particle Formation in the South Aegean Sea during the Etesians: importance for CCN production and cloud droplet number” by P. Kalkavouras et al.

P. Kalkavouras et al.

mtombrou@phys.uoa.gr

Received and published: 14 October 2016

We would like to thank both reviewers for their comments and recommendations. We believe that we have corrected and improved the paper by incorporating their comments, in the revised version. The figure proposed by the first reviewer was a very good idea where we had to clarify several points of ‘our story’ to provide sufficient context. We reran the simulations at higher resolution, replaced figures and modified the discussion, accordingly. The main changes are the following: Following both Reviewers’ comment, regarding the model’s estimation of the simulated new particle formation, we reran the model by ignoring NPF process. In the revised manuscript, section 3.5 is divided in 2 sections: 3.5 is called “Impact of NPF events on CCN production”

C1

and 3.6 “Impact of NPF events on cloud droplet number.” We followed first reviewer’s suggestion to use for the two types of northern flow the terms: Etesian Flow (EF) and Moderate Surface Flow (MSF), in order to have a more concise wording. We also followed the same formalism in the revised Tables and Figure captions.

Reviewer #1:

General Comments: Kalkavouras and co-authors present results from an intriguing experiment in the Aegean Sea. The nature of pollution arriving at the long-standing Finokalia measurement platform is investigated directly with observations at Santorini, a site strategically located along the trajectory from the European mainland. The observations target aerosol size distributions, relevant primary and secondary pollutant concentrations, and detection of new particle formation events. The authors identify two events representing characteristic flow from the north, albeit of two distinct types. The Etesian flow example is marked with NPF events at both sites, although the events are stronger at the Santorini site. The authors extend their observations of particle number and composition to predictions of CCN at various supersaturations. They also go further to predict the total effect of the NPF events on cloud droplet number, taking into account the impact of constrained available water vapor. The paper starts with a nice scientific idea and goes into good detail into the results. What I see lacking is a little more connective tissue linking the observations at each site with each other and the mainland into a cohesive story. The material is already there, but it is somewhat buried and could be highlighted with a figure, for instance. I would like to see the following points addressed before recommending publication:

Specific Comments: 1. I recommend the authors adopt two shorthand names for the distinct periods (22/7- 24/7 and 25/7-27/7). They could be referred to as “Etesian Flow” and “Moderate Surface Flow,” for example. Small changes like this could help the readability of the paper. A useful addition to this paper would be a two-panel cartoon, each overlaid on the Greek domain map in Fig. 1, for example, that describes the factors at play in these two periods. They could identify generally where they expect

C2

emissions, mixing, oxidation, NPF, and aging of new particles to be happening.

To be more concise, we used for the two types of northern flow the terms: Etesian Flow (EF) and Moderate Surface Flow (MSF). We also followed the same formalism in the revised Tables and Figure captions. Figure 1 is replaced with a two-panel cartoon in order to identify the locations where we expect/identify the various processes.

2. The WRF-Chem aerosol module configuration, as documented by the authors, is problematic for this particular application. It is quite likely that NPF events and subsequent processing are not captured realistically at all by the model. The sulfuric acid/water pathway parameterized by Kulmala et al. (1998) is likely not strong enough to enhance particles near the surface and lower troposphere to the levels observed at the Santorini site. It is now well-documented that other reagents play important roles in this process (e.g. NH₃ and organics), and these pollutants have been identified by the authors to be present and significant components of the aerosol. My guess is that most of the Aitken-mode particles in the model originate from direct emissions, not from secondary generation. A related issue is the lack of a dedicated nucleation mode in the model. Without this mode, any NPF events will artificially broaden the Aitken mode distribution and give unrealistic lifetimes against deposition and coagulation. It will also affect the growth rates predicted by the model. The authors astutely sidestep relying on the model to predict size distributions and use their own observations when possible for calculating CCN and cloud drop number concentrations. However, since they include an entire section (2.2) detailing the regional modeling they performed, it is a good idea to explicitly state the limitations of this analysis for particle size distributions, and remind the reader that they are using the WRF-Chem output for its knowledge of advection flows and chemical composition, not microphysical processing.

We agree with the Reviewer's comment. In version 3.3 the aerosol models are not appropriate to simulate the NPF events realistically. Luo and Yu (2011), discuss the need to improve the representation of the nucleation process in earlier versions of WRF-Chem. We believe that more research is needed regarding the nucleation modeling

C3

in the area, which we plan to perform in a separate paper in the near future. Nevertheless, we conducted another simulation ignoring the nucleation parameterization in order to comprehend/emphasize the spatial extent of these processes and present them in Fig. 1. In the revised manuscript, the relevant discussion on model limitations (due to lack of a dedicated nucleation mode, nucleation parameterization) is presented in section 2.3 'Regional modeling' (page 7, lines 6-21).

To further elucidate the conditions under which NPF events take place in our region, we reran the model by ignoring the NPF process. In the revised manuscript our hypothesis is not based on the simulated Aitken-mode particles but on the number concentration differences considering and ignoring NPF process. The relevant discussion is presented in section 3.4 'Spatial extent of NPF event' (pages 13-16).

3. The authors conclude that the NPF events observed at Santorini are regional in nature with a spatial scale of 250 km and characteristic transport time of 4.5 hours. They also assert that Finokalia does not see the bursts because it is 3 hours away and particle populations age before they arrive there. The authors do note that this second observation demonstrates how site-to-site variability can be important during a regional event. I am not sure that this totally addresses the issue though. Why are the events sort of regional and sort of not-regional? Is this an issue of using up the NPF precursors before the air mass gets to Santorini and then shifting to chemical conditions that favor condensation to available surface area? Or is something else at play here?

We used the regional characterization, mainly because the number of particles remained high for several hours at Santorini (Kulmala et al., 2012). In addition, the NPF event was found to extend over hundreds of kilometers. Thus, as the reviewer points out, we tried to relate these fine aerosols to regional sources of pollution transferred by long-range transport (LRT) during Etesian flow conditions. Despite that, we observed local variability at sub-regional scales, due to the differences in geographical and atmospheric conditions between stations along the same trajectory. This is the

C4

case between Santorini and Finokalia stations. We also expect that local variability is unavoidable at smaller scales, over locations at the same distance from the sources. This is based on the simulations that show that the spatial differences of chemical and physical properties in the initial steps of the formation, under the stable Etesian flow, produce streams with different characteristics, especially upwind of Santorini.

4. The paragraph beginning line 13 on page 9 describes an interesting hypothesis for how pollutants are transported to the middle of the Aegean Sea with limited aging. However, I'm not convinced there is enough evidence to warrant the detailed discussion that is given to this possibility or the certainty with which it is treated in the conclusions section. As described in my first comment, any model data related to the size distribution of Aitken-mode particles probably cannot be trusted in this case. If I understand correctly, the main assertion here is that the particles were formed over the Turkish mainland and transported quickly before they have a chance to be significantly coagulated away. Why could the enhanced number concentrations not come instead from oxidation and NPF over the water during transport, where there may be enhanced photochemistry, complex interactions with clouds, interesting boundary layer phenomena, etc? If I'm not understanding the meat of the argument correctly, please explain it to me and consider rephrasing it in the text to be clearer. What insight do the model CO concentrations help to provide regarding the stratification and mixing of distinct layers downwind of the continent?

The revised section 3.4 'Spatial extent of NPF event' (pages 13-16) explains more clearly now the processes taking place. The discussion is mainly based on the number concentration differences considering and ignoring NPF process and not on the Aitken-mode simulated particles. Although the model severely underestimates the NPF, the decisive role of the Etesian flow on the evolution of the phenomenon over the Aegean Sea is evident (from page 13 line 24 to page 15 line 1). The atmospheric conditions under a similar Etesian event have been studied thoroughly in a separate paper that has been submitted to BLM. In particular, the heat fluxes simulated and calculated

C5

from airborne measurements over the AS (Tombrou et al., 2015) varied from -25 W m⁻² (over the northeastern AS) to 25 W m⁻² (over the southeastern AS). Furthermore, vertical cross-sections of measured CO concentrations along the eastern AS under an Etesian flow, are shown in Fig. 7 by Tombrou et al. (2015). The strong gradient of stratification and mixing downwind of AS, is apparent. In particular, at 40° latitude, where the plume leaves the Turkey continent, the vertical mixing extends up to 500-600m height according to the CO vertical extent. Above the Cyclades complex (lat 36.5° - 38°) the mixing extends up to 1km and gradually increases up to 2km, upwind of Crete (Finokalia at 35°).

5. I recommend separating the paragraphs detailing the Nd calculations (starting on Page 12) into their own section, perhaps called "Impact of NPF events on cloud droplet number." Then section 3.5 would be called "Impact of NPF events on CCN production."

Done

6. How is the partial sensitivity of cloud droplet number to chemical composition and vertical velocity determined? Can the equations be provided? What is the uncertainty associated with this? Please document it if possible.

The reviewer raises a good question. The sensitivity is derived from the parameterization using either a direct sensitivity or finite difference approach, as described in Karydis et al. (2012). Here we use the finite difference implementation. This information is now given on page 6; Ln 16-18. The accuracy of the method, i.e. the ability of the parameterization to capture the sensitivity of droplet number to each parameter examined was explored in detail by Morales and Nenes (2014). Given that the parameterization has been shown to give cloud droplet closure in ambient clouds to within experimental uncertainty (Meskhidze et al., 2005; Fountoukis et al., 2007; Hoyle et al., 2016), and that the same parameterization also reproduces the droplet number and sensitivities of the detailed numerical simulation with high fidelity (Morales and Nenes, 2014), we expect the sensitivities and attribution calculations presented here to be

C6

representative of the ambient clouds in the study region.

Minor Changes/Typos: Pg 2, Ln 25: The phrase “without any particular seasonal preference about their occurrence” is difficult to understand. Can the authors please reword this to be more specific?

The phrase has been replaced by (Pg 3, Ln 11-12): “Most of these ground-based observations indicate that the mass of fine aerosols presents a summer maximum, however the frequency of the events is season independent.”

Pg 3, Ln 4: “prior to reaching”

Done

Pg 5, Ln 11-13: This is technically not a sentence.

Replaced by (page 8, lines 6-8): “Air mass origin and trajectories were determined by HYSPLIT4 (Hybrid Single-Particle Lagrangian Integrated Trajectory; www.arl.noaa.gov/ready/hysplit4.html) back-trajectory analysis (Draxler and Rolph, 2015).”

Fig. 3 and caption: “open circles” not “cycles”.

Done.

Also, please indicate on the figure that the solid lines describe wind speed and the circles describe direction. It is hinted at in the figure and explained in the caption, but it would be quicker for the reader to have it identified visually, with an arrow or something.

Figure 3 is replaced by a new one, where the abbreviations ‘ws’ and ‘wd’ are now included indicating wind speed and wind direction respectively. The caption is rephrased accordingly: ‘Time series of the wind speeds (ws, solid lines on left axis) and wind directions (wd, open circles on right axis) at Santorini (simulations by the WRF-Chem model) and at Finokalia (measurements). The second period of the EF is shaded with yellow and the MSF period with grey.’

C7

Pg 6, Ln 16-17: The “less pronounced” diurnal cycle at Finokalia for ozone is not obvious to me from Fig. S4. Please include a plot of the actual diurnally averaged profiles or report the daily minima and maxima to demonstrate this point.

We refer to the Etesian period (EF) that corresponds to the yellow panel in Fig. S4. The mean diurnal range at Finokalia station (from 21 to 24 July) is 8 ppbv, while at Santorini, for the same period is 18 ppbv (Fig. S4). This information is now included in the text (page 9, lines 22-24)

Pg 6, Ln 28-29: I would not characterize -21% or -15% under-prediction as “small”. Either establish what they are small compared to, or please get rid of this qualification.

We agree with the reviewer’s comment, therefore, we decided to delete the word ‘small’.

Pg 6, Ln27-30: Please break this sentence up. It is long and confusing.

Replaced by (page 10, lines 9-13): Simulations confirm that the air masses received at both stations during the prevailing strong northern wind are of the same origin, and representative of EF conditions (Fig. S3) albeit with an O₃ under-prediction (average bias during afternoon hours up to -21% on 23 and -15% on 24 July, Fig. 5). During the MSF period, simulations indicate an O₃ increase, especially in the southern AS, but also underpredicted (average bias during afternoon hours up to -24% on 26 July, Fig. 5).

Pg 6, Ln 31-32: Is there a more recent or relevant reference than McKeen et al. (1991)? Anything that specifically identifies this model scenario or modern European scenarios in general as suffering from ozone boundary condition issues?

The chemical boundary conditions used in this modeling study are hardcoded in the WRF-Chem model. The values are based on an idealized, northern hemispheric, mid-latitude, clean environmental, vertical profile from the NOAA Aeronomy Lab Regional Oxidant Model (NALROM) (Liu et al. 1996; Peckham et al. 2011). This information has been added to section 2.3 ‘Regional modeling’ (page 7, lines 25-27 while the reference

C8

of McKeen et al. (1991) was omitted.

Pg 7, Ln 5-6: In what way did the inorganic and organic mass concentrations show “similar behavior” to that of ozone? Are the authors just identifying them all as secondary pollutants? Please provide an estimate of the correlation coefficient or index of agreement for a statement like this.

We identify all of them as secondary pollutants driven by the same meteorological conditions. The R2 of O3 to Organics and O3 to inorganics is 0.5 and 0.59 respectively (included in the revised manuscript on page 10, line 22). Also, the simulated spatial patterns of O3 and sulfates are similar, for each period (EF and MSF).

Pg 7, Ln 9: Please remove the comma after the parentheses.

Done

Pg 7, Ln 20: Please refer to some quantitative statistics to back up this claim.

An extended evaluation of WRF-Chem model against airborne and ground observations over the AS during the Etesians is presented in Bossioli et al. (2016). In that study biomass burning emissions were also included. After the reviewer’s suggestion some statistics have been added in the revised manuscript: For EF period (Page 11, lines 3-4): “...on average during EF underprediction of 30% for sulfates and 60% for ammonium” For MSF period (Page 11, lines 12-15): “(simulated and observed concentrations correlate during both periods R2=0.8), however they are lower than the measured values at Finokalia station (on average underprediction of 50% for sulfates and 75% for ammonium).

Pg 10, Ln 6-8: This sentence is worded in a confusing way.

The discussion has been revised (page 15, lines 10-17). The specific sentence has been revised to “The nucleation-mode particles are significantly reduced as they have shifted gradually towards larger sizes (Aitken-mode), before reaching Finokalia (Fig. 4).” Pg 10, Ln 25 – Pg 11, Ln 19: Most of this material would be better-placed in the

C9

methods section (2.3 maybe). This goes for the second paragraph a page 12 as well.

Done

References:

Bossioli, E., Tombrou, M., Kalogiros, J., Allan, J., Bacak, A., Bezantakos, S., Biskos, G., Coe, H., Jones, B. T., Kouvarakis, G., Mihalopoulos, N., and Percival, C. J.: Atmospheric composition in the Eastern Mediterranean: Influence of biomass burning during summertime using the WRF-Chem model, *Atmos. Environ.*, 132, 317–331, 2016

Draxler, R.R. and Rolph, G. D.: HYSPLIT (HYbrid Single-Particle Lagrangian Integrated Trajectory) Model access via NOAA ARL READY Website (<http://ready.arl.noaa.gov/HYSPLIT.php>). NOAA Air Resources Laboratory, Silver Spring, MD., 2015

Fountoukis, C., Nenes, A., Meskhidze, N., Bahreini, R., Brechtel, F., Conant, W. C., Jonsson, H., Murphy, S., Sorooshian, A., Varutbangkul, V., R. C. Flagan, and J. H. Seinfeld Aerosol–cloud drop concentration closure for clouds sampled during ICARTT, *J.Geoph.Res.*, 112, D10S30, doi:10.1029/2006JD007272, 2007 Hoyle, C.R., Webster, C.S., Rieder, H.E., Nenes, A., Hammer, E., Herrmann, E., Gysel, M., Bukowiecki, N., Weingartner, E., Steinbacher, M., and U. Baltensperger Chemical and physical influences on aerosol activation in liquid clouds: a study based on observations from the Jungfrauoch, Switzerland, *Atmos.Chem.Phys.*, 16, 4043–4061, 2016 Karydis, V. A., Capps, S. L., Russell, A. G., and Nenes, A.: Adjoint sensitivity of global cloud droplet number to aerosol and dynamical parameters, *Atmos. Chem. Phys.*, 12, 9041-9055, doi:10.5194/acp-12-9041-2012, 2012 Kulmala, M., Petäjä, T., Nieminen, T., Sipilä, M., Manninen, H. E., Lehtipalo, K., Dal Maso, M., Aalto, P. P., Junninen, H., Paasonen, P., Riipinen, I., Lehtinen, K. E. J., Laaksonen, A., and Kerminen, V.-M.: Measurement of the nucleation of atmospheric aerosol particles. *Nature Protocols* 7, 1651–1667, 2012.

Liu, S. C., McKeen, S. A., Hsie, ÎŤ-Îě., Lin, ., Kelly, K. K., Bradshaw, J. D., Sand-

C10

holm, S. T., Browell, E. V., Gregory, G. L., Sachse, G. W., Bandy, A. R., Thornton, D. C., Blake, D. R., Rowland, F. S., Newell, R., Heikes, B. G., Singh, H., Talbot, R. W.: Model study of tropospheric trace species distributions during PEM-West A, *J. Geophys. Res.*, 101(D1), 2073–2085, doi:10.1029/95JD02277, 1996 Luo, G. and Yu, F.: Simulation of particle formation and number concentration over the Eastern United States with the WRF-Chem + APM model, *Atmos. Chem. Phys.*, 11, 11521–11533, doi:10.5194/acp-11-11521-2011, 2011. Meskhidze, N., A. Nenes, Conant, W. C., and Seinfeld, J.H.: Evaluation of a new Cloud Droplet Activation Parameterization with In Situ Data from CRYSTAL-FACE and CSTRIFE, *J. Geophys. Res.*, 110, D16202, doi:10.1029/2004JD005703, 2005 Morales Betancourt, R., and Nenes, A.: Aerosol Activation Parameterization: The population splitting concept revisited, *Geosci. Mod. Dev.*, 7, 2345–2357, 2014 Peckham, S., G. A. Grell, S. A. McKeen, M. Barth, G. Pfister, C. Wiedinmyer, J. D. Fast, W. I. Gustafson, R. Zaveri, R. C. Easter, J. Barnard, E. Chapman, M. Hewson, R. Schmitz, M. Salzmann, S. Freitas: WRF/Chem Version 3.3 User's Guide. NOAA Technical Memo., 98 pp., 2011

Tombrou, M., Bossioli, E., Kalogiros, J., Allan, J. D., Bacak, A., Biskos, G., Coe, H., Dandou, A., Kouvarakis, G., Mihalopoulos, N., Percival, C. J., Protonotariou, A. P., and Szabó-Takács, B.: Physical and chemical processes of air masses in the Aegean Sea during Etesians: Aegean-GAME airborne campaign, *Sci. Total Environ.*, 506–507, 201–216, doi:10.1016/j.scitotenv.2014.10.098, 2015

Interactive comment on *Atmos. Chem. Phys. Discuss.*, doi:10.5194/acp-2016-330, 2016.