



#### 1 Influence of Seed Aerosol Surface Area and Oxidation Rate on Vapor-Wall 2 Deposition and SOA Mass Yields: A case study with α-pinene Ozonolysis

3 T. Nah,<sup>1</sup> R. C. McVay,<sup>2</sup> X. Zhang,<sup>3,#</sup> C. M. Boyd,<sup>1</sup> J. H. Seinfeld<sup>2,3</sup> and N. L. Ng<sup>1,4\*</sup> 4

5 6 7 <sup>1</sup>School of Chemical and Biomolecular Engineering, Georgia Institute of Technology, Atlanta, GA, USA

<sup>2</sup>Division of Chemistry and Chemical Engineering, California Institute of Technology, Pasadena, CA, USA

8 9 <sup>4</sup>School of Earth and Atmospheric Sciences, Georgia Institute of Technology, Atlanta, GA, USA

11 12

\* To whom correspondence should be addressed: ng@chbe.gatech.edu

- 13
- 14

### Abstract

15 Laboratory chambers, invaluable in atmospheric chemistry and aerosol formation studies, 16 are subject to particle and vapor wall deposition, processes that need to be accounted for 17 in order to accurately determine secondary organic aerosol (SOA) mass yields. Although particle wall deposition is reasonably well understood and usually accounted for, vapor 18 19 wall deposition is less so. The effects of vapor wall deposition on SOA mass yields in 20 chamber experiments can be constrained experimentally by increasing the seed aerosol 21 surface area to promote the preferential condensation of SOA-forming vapors onto seed 22 aerosol. Here, we study the influence of seed aerosol surface area and oxidation rate on 23 SOA formation in  $\alpha$ -pinene ozonolysis. The observations are analyzed using a coupled 24 vapor-particle dynamics model to interpret the roles of gas-particle partitioning (quasi-25 equilibrium vs. kinetically-limited SOA growth) and  $\alpha$ -pinene oxidation rate in 26 influencing vapor wall deposition. We find that the SOA growth rate and mass yields are 27 independent of seed surface area within the range of seed surface area concentrations 28 used in this study. This behavior arises when the condensation of SOA-forming vapors is 29 dominated by quasi-equilibrium growth. Faster  $\alpha$ -pinene oxidation rates and higher SOA 30 mass yields are observed at increasing  $O_3$  concentrations for the same initial  $\alpha$ -pinene 31 concentration. When the  $\alpha$ -pinene oxidation rate increases relative to vapor wall 32 deposition, rapidly produced SOA-forming oxidation products condense more readily 33 onto seed aerosol particles, resulting in higher SOA mass yields. Our results indicate that 34 the extent to which vapor wall deposition affects SOA mass yields depends on the

<sup>&</sup>lt;sup>3</sup>Division of Engineering and Applied Science, California Institute of Technology, Pasadena, CA, USA

<sup>10</sup> <sup>#</sup>Now at: Center for Aerosol and Cloud Chemistry, Aerodyne Research, Billerica, MA, USA





35 particular VOC system, and can be mitigated through the use of excess oxidant

36 concentrations.

# 37 **1. Introduction**

38 Secondary organic aerosol (SOA), formed from the oxidation of volatile and 39 intermediate volatility organic compounds (VOCs and IVOCs), contributes a significant 40 fraction of the global organic aerosol burden (Kanakidou et al., 2005; Hallquist et al., 2009; Tsigaridis et al., 2014). SOA formation studies, which are typically conducted in 41 42 laboratory chambers in the presence of seed aerosol particles, provide fundamental data 43 that can be used to predict the rate of atmospheric SOA formation. An essential 44 parameter of interest in laboratory chamber studies is the SOA mass yield (Y), which is 45 defined as the ratio of mass concentration of SOA formed to mass concentration of parent hydrocarbon reacted (( $\Delta$ HC), Y =  $\Delta$ M<sub>0</sub>/ $\Delta$ HC (Odum et al., 1996; Odum et al., 1997a; 46 47 Odum et al., 1997b)). The measured SOA mass yields can subsequently be applied in 48 atmospheric models to predict regional and global organic aerosol burdens. In order to 49 obtain accurate SOA mass yields from the evolving aerosol size distribution in chamber 50 experiments, the loss of both particles and vapors to the chamber walls needs to be 51 accurately accounted for (Crump and Seinfeld, 1981; McMurry and Grosjean, 1985; McMurry and Rader, 1985; Cocker et al., 2001a; Weitkamp et al., 2007; Pierce et al., 52 53 2008; Hildebrandt et al., 2009; Loza et al., 2010; Matsunaga and Ziemann, 2010; Loza et al., 2012; Kokkola et al., 2014; McVay et al., 2014; Yeh and Ziemann, 2014; Zhang et 54 55 al., 2014; Yeh and Ziemann, 2015; Zhang et al., 2015a; La et al., 2016).

56 The mechanisms by which particles in chambers deposit on chamber walls are 57 reasonably well understood. Particles are transported to the boundary layer on the 58 chamber walls via diffusion, gravitational settling, and electrostatic forces (Crump and 59 Seinfeld, 1981; McMurry and Grosjean, 1985; McMurry and Rader, 1985; Pierce et al., 2008). The rate at which particles are transported to the edge of the boundary layer is 60 61 dictated primarily by mixing conditions in the chamber. An effective approach for 62 characterizing particle wall loss involves measuring the size-dependent wall loss rates of polydisperse inert seed aerosol (e.g. ammonium sulfate particles) injected into the 63 chamber during seed-only experiments (Keywood et al., 2004; Pierce et al., 2008). The 64





65 observed particle number concentration decay in each size bin is then fitted to a first-66 order exponential decay from which the first-order wall loss coefficients are determined 67 as a function of particle size. These wall loss coefficients are subsequently used to correct 68 for size-dependent particle wall loss in actual SOA formation experiments.

69 Vapor-wall deposition mechanisms in chambers are not as well understood or 70 accounted for as those for particles. The degree to which SOA-forming vapors deposit 71 onto chamber walls is governed by the rate at which these gas-phase organic molecules 72 are transported to the walls, the strength of adherence of the organic molecule to the wall, 73 and the extent of reversible vapor-wall partitioning (Loza et al., 2010; Matsunaga and 74 Ziemann, 2010; Zhang et al., 2015a). For example, Loza et al. (2010) showed that the 75 loss of 2,3-epoxy-1,4-butanediol, an isoprene oxidation product analogue, to walls in the 76 Caltech chamber was essentially irreversible on short time scales but became reversible 77 on longer time scales. In contrast, glyoxal, a common isoprene oxidation product, 78 exhibited reversible vapor-wall partitioning over all time scales. Recent studies show that 79 SOA mass yields measured in chamber experiments can be significantly underestimated 80 due to wall deposition of SOA-forming vapors that would otherwise contribute to SOA 81 growth (McVay et al., 2014; Zhang et al., 2014; La et al., 2016). Zhang et al. (2014) 82 found that chamber-derived SOA mass yields from toluene photooxidation may be 83 underestimated by as much as a factor of four as a result of vapor wall loss. 84 Consequently, the use of underestimated chamber-derived SOA mass yields in 85 atmospheric models will lead to the underprediction of ambient SOA mass concentrations 86 (Cappa et al., 2016).

87 For the toluene photooxidation system, Zhang et al. (2014) showed that the measured SOA mass yields increased with increasing seed aerosol surface area, 88 89 demonstrating that increasing the seed-to-chamber surface area ratio promoted the 90 condensation of SOA-forming vapors onto seed aerosol particles. However, increasing 91 the seed aerosol surface area to promote condensation of SOA-forming vapors onto seed aerosol particles may not be effective in all VOC oxidation systems. A modeling study by 92 93 McVay et al. (2014) showed that the SOA mass yield depends on seed aerosol surface 94 area only in cases where the condensation of SOA-forming vapors onto seed aerosol





95 particles is kinetically limited (i.e., the timescale for gas-particle equilibrium is 96 competitive with or greater than the timescale for reaction and vapor-wall deposition). In 97 addition to the seed aerosol surface area, VOC oxidation rate may also play an important 98 role in the effect of vapor wall loss on SOA formation. Ng et al. (2007) showed that the 99 SOA mass yields from *m*-xylene photooxidation are dependent on the oxidation rate, with 100 higher OH concentrations (hence faster oxidation rates) resulting in higher SOA mass 101 yields. It was suggested that the "oxidation rate effect" could arise as a result of 102 competition between growing particles and chamber walls for condensable VOC 103 oxidation products (Ng et al., 2007). However, McVay et al. (2016) reported similar SOA 104 growth at low and high OH concentrations in α-pinene photooxidation. Taken together, 105 these studies show the importance of understanding how gas-particle partitioning and 106 VOC oxidation rate impact vapor-wall deposition and SOA mass yields in laboratory 107 chamber experiments.

108 In this study, we examine the influence of seed aerosol surface area and oxidation 109 rate on SOA formation in α-pinene ozonolysis chamber experiments. α-pinene is the most abundant monoterpene, with global emissions estimated to be  $\sim 66 \text{ Tg yr}^{-1}$  (Guenther et 110 111 al., 2012). Ozonolysis is the major atmospheric oxidation pathway of  $\alpha$ -pinene, and is 112 estimated to account for reaction of ~46 % of emitted  $\alpha$ -pinene (Griffin et al., 1999; 113 Capouet et al., 2008). α-pinene ozonolysis, a major source of atmospheric SOA on both 114 regional and global scales (Kanakidou et al., 2005; Hallquist et al., 2009; Carlton et al., 2010; Pve et al., 2010), has been the subject of numerous studies (Hoffmann et al., 1997; 115 116 Griffin et al., 1999; Cocker et al., 2001b; Gao et al., 2004; Presto et al., 2005; Presto and Donahue, 2006; Pathak et al., 2007a; Pathak et al., 2007b; Song et al., 2007; Shilling et 117 118 al., 2008; Henry et al., 2012; Ehn et al., 2014; Kristensen et al., 2014; Zhang et al., 119 2015b). Here, we measure the  $\alpha$ -pinene SOA mass yield as a function of seed aerosol surface area concentration (0 to 3000  $\mu$ m<sup>2</sup> cm<sup>-3</sup>) and O<sub>3</sub> mixing ratio (100 vs. 500 ppb). 120 121 These results are analyzed using a coupled vapor-particle dynamics model to evaluate the 122 roles of gas-particle partitioning and VOC oxidation rate in influencing vapor-wall 123 deposition effects on the measured SOA mass yields.

### 124 **2. Experimental**





#### 125 **2.1. Dark** α-pinene ozonolysis experiments

Experiments were conducted in the Georgia Tech Environmental Chamber (GTEC) facility. Details of the dual chamber facility are provided elsewhere (Boyd et al., 2015). Only one FEP Teflon chamber (volume 13 m<sup>3</sup>) was used for the entirety of this study. Before each experiment, the chamber is flushed with dried, purified air for at least 36 h until the aerosol number concentration is < 30 cm<sup>-3</sup>. All experiments were conducted under dry conditions (< 5 % RH) at room temperature (25 °C). NO<sub>x</sub> mixing ratios in these experiments are < 1 ppb. Experimental conditions are summarized in Table 1.

133 22 ppm of cyclohexane (Sigma Aldrich,  $\geq$  99.9 %) was first injected into the 134 chamber to act as an OH scavenger (~440 times the initial  $\alpha$ -pinene concentration). After 135 the cyclohexane concentration had stabilized in the chamber for 30 min, a known concentration (~50 ppb in all experiments) of  $\alpha$ -pinene (Sigma Aldrich, > 99 %) was 136 injected into the chamber, followed by inorganic seed aerosol via atomization of an 137 138 aqueous ammonium sulfate (AS) solution (in seeded experiments). To vary the seed 139 aerosol surface area, different concentrations of AS solutions were used to generate seed 140 aerosol particles in the seeded experiments. In the "low AS" experiments, a 0.015 M AS solution was used to generate seed particles, and the resulting initial total AS seed surface 141 area concentration was  $\sim 1000 \text{ }\mu\text{m}^2 \text{ cm}^{-3}$ . In the "high AS" experiments, a 0.05 M AS 142 143 solution was used to generate seed aerosol particles, and the resulting initial total AS seed surface area concentration was  $\sim 3000 \text{ } \text{ } \text{ } \text{m}^{-3}$ . In selected experiments, no seed aerosol 144 145 particles were introduced into the chamber and SOA was formed via nucleation. After the 146 seed aerosol concentration in the chamber stabilized,  $O_3$  (100 or 500 ppb), which was 147 generated by passing purified air into a photochemical cell (Jelight 610), was introduced 148 into the chamber. The start of  $O_3$  injection into the chamber marked the beginning of the 149 reaction (i.e., reaction time = 0 min). The initial  $\alpha$ -pinene:O<sub>3</sub> molar ratio was fixed at approximately 1:2 and 1:10 in the 100 and 500 ppb O<sub>3</sub> experiments, respectively. O<sub>3</sub> was 150 151 injected into the chamber for 13.5 and 54.25 min in the 100 and 500 ppb O<sub>3</sub> experiments, 152 respectively, to achieve the desired  $O_3$  concentrations. The  $O_3$  injection times were 153 established in separate experiments in which only O<sub>3</sub> was injected into the chamber.





154 The  $\alpha$ -pinene and O<sub>3</sub> concentrations were measured by a Gas Chromatograph-155 Flame Ionization Detector (GC-FID, Agilent 7890A) and O<sub>3</sub> monitor (Teledyne T400), 156 respectively. GC-FID measurements were taken 12 min apart. A High Resolution Time-157 of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS, Aerodyne Research Inc.) was used 158 to measure the aerosol elemental composition (DeCarlo et al., 2006; Canagaratna et al., 159 2015). Details on the operation of the HR-ToF-AMS and its data analysis are described 160 elsewhere (Canagaratna et al., 2015). Aerosol size distributions, number and volume 161 concentrations were measured by a Scanning Mobility Particle Sizer (SMPS, TSI), which 162 consists of a Differential Mobility Analyzer (DMA, TSI 3081) and a Condensation 163 Particle Counter (CPC, TSI 3775). For nucleation and low AS experiments, the measured 164 aerosol size range was set to 14 to 686 nm diameter. For high AS experiments, the 165 measured aerosol size range was set to 17 to 983 nm. Prior checks were made to confirm 166 that no particles larger than 686 nm were detected in the nucleation and low AS 167 experiments. The SOA mass concentrations reported in this study were measured using 168 the SMPS. The SOA density was calculated from the ratio of the aerosol size 169 distributions measured by the HR-ToF-AMS and the SMPS during nucleation 170 experiments (DeCarlo et al., 2004; Bahreini et al., 2005).

171 **2.2. Particle wall deposition correction** 

172 Particle wall deposition needs to be accounted for to determine the SOA mass 173 concentration in the chamber. Two limiting assumptions have traditionally been made 174 regarding interactions between particles deposited on the chamber walls and suspended 175 vapors when accounting for particle wall loss in the computation of SOA mass yields 176 (Weitkamp et al., 2007; Hildebrandt et al., 2009; Loza et al., 2012; Zhang et al., 2014). 177 The first case assumes that particles deposited on the walls cease to interact with 178 suspended vapors, and therefore the SOA mass present on these deposited particles does 179 not change after deposition (Loza et al., 2012; Zhang et al., 2014). Adding the SOA mass 180 present on these deposited particles to that present on the suspended particles provides a 181 lower bound of the total SOA mass concentration. In the second case, it is assumed that 182 particles deposited on the walls continue to interact with suspended vapors as if these particles had remained suspended, and therefore the SOA mass present on these 183





184 deposited particles increases at the same rate as those suspended (Hildebrandt et al., 185 2009; Weitkamp et al., 2007). Thus, this case provides an upper bound of the total SOA 186 mass concentration due to the additional uptake of suspended vapors to wall-deposited 187 particles. However, it must be kept in mind that the calculated SOA mass concentration 188 can be underestimated even in the upper bound case since the calculation accounts 189 neither for differences in the vapor-particle and vapor-wall interaction and transport 190 timescales nor for the significantly larger amount of absorbing mass of the chamber walls 191 (relative to the deposited particles) for suspended vapors (McVay et al., 2014; Zhang et 192 al., 2014; McVay et al., 2016).

In this study, we calculate SOA mass yields using the lower bound of the total SOA mass concentration obtained from SMPS measurements, which has been described in detail previously (Loza et al., 2012), and will be reviewed briefly here. For each particle size bin *i* at each time increment  $\Delta t$ , the particle number distribution deposited on the wall  $(n_{w,i,j})$  is:

198 
$$n_{w,i,j} = n_{s,i,j} \times (1 - \exp(-\beta_i \Delta t))$$
(1)

where  $n_{s,i,j}$  is the suspended particle number distribution in particle size bin *i* at time step *j*,  $\Delta t$  is the difference between time step *j* and time step *j* + 1, and  $\beta_i$  is the size-dependent first-order exponential wall loss rate obtained from seed-only experiments. The particle wall loss corrected number distribution ( $n_{total,i,j}$ ) is obtained from the sum of the particle number distribution of deposited particles ( $n_{w,i,j}$ ) and suspended particles ( $n_{s,i,j}$ ):

204 
$$n_{total,i,j} = n_{s,i,j} + n_{w,i,j}$$
(2)

Assuming spherical particles, the particle wall loss corrected volume concentration  $(V_{total,j})$  is:

207 
$$V_{total,j} = \sum_{i=1}^{m} \frac{n_{total,i,j}}{D_{p,i} \ln 10} \times (D_{p,i+} - D_{p,i-}) \times \frac{\pi}{6} D_{p,i}^{3}$$
(3)

where *m* is the number of particle size bins,  $D_{p,i^+}$  and  $D_{p,i^-}$  are the upper and lower limits for size bin *i*, respectively, and  $D_{p,i}$  is the median particle diameter for size bin *i*. The term





210  $D_{p,i}$ ln10 is needed to convert from a lognormal distribution. Figures S1-S4 and Table S1 211 show results from the particle wall loss correction. To calculate the SOA mass 212 concentration ( $\Delta M_{o,j}$ ), the SOA density ( $\rho_{org}$ ) is multiplied by the difference of the 213 particle wall loss corrected volume concentration ( $V_{total,j}$ ) and the initial seed volume 214 concentration ( $V_{seed}$ ):

$$\Delta M_{o,i} = \rho_{org} \times (V_{total,i} - V_{seed})$$
<sup>(4)</sup>

The measured densities of the  $\alpha$ -pinene SOA are 1.39 and 1.37 g cm<sup>-3</sup> for the 100 and 500 ppb O<sub>3</sub> experiments, respectively, and are within the range (i.e., 1.19 to 1.52 g cm<sup>-3</sup>) reported in previous  $\alpha$ -pinene ozonolysis studies (Bahreini et al., 2005; Kostenidou et al., 2007; Song et al., 2007; Shilling et al., 2009).

### 220 **3. Vapor-particle dynamics model**

221 A coupled vapor-particle dynamics model is used to evaluate the influence of seed 222 aerosol surface area and oxidation rate on SOA formation in the  $\alpha$ -pinene ozonolysis 223 chamber experiments. This model is similar to that used in McVay et al. (2014), and will 224 be briefly described here. Parameters from the experimental data (temperature, pressure, 225 initial  $\alpha$ -pinene concentration) are used as model inputs. The initial size distribution is set 226 to that measured by the SMPS, with the exception of the two nucleation experiments. 227 Because nucleation is not explicitly simulated, an approximation is used in which the 228 smallest diameter bin is initialized with the total number of particles measured at the end 229 of the experiment (see Table S1). In each simulation, the decay of  $\alpha$ -pinene, the consumption of O<sub>3</sub>, the SOA mass concentration, and the SOA mass yield are calculated 230 throughout the duration of the experiment. We assume a linear injection rate of O<sub>3</sub> based 231 232 on the time required to inject the desired O<sub>3</sub> concentration. For example, O<sub>3</sub> is injected at a rate of 500/54.25 ppb min<sup>-1</sup> for the first 54.25 min during the 500 ppb  $O_3$  experiments. 233  $O_3$  simultaneously decays by reaction with  $\alpha$ -pinene at a rate constant of  $9.4 \times 10^{-17}$  cm<sup>3</sup> 234 molec.<sup>-1</sup> s<sup>-1</sup> (Saunders et al., 2003). The  $O_3+\alpha$ -pinene reaction is assumed to occur in a 235 well-mixed chamber and produces 5 classes of first-generation products, which are 236 grouped according to mass saturation concentrations, similar to the volatility basis set 237 (Donahue et al., 2006): >10<sup>3</sup> (assumed to be completely volatile),  $10^2$ , 10, 1 and 0.1 µg 238





m<sup>-3</sup>. Branching ratios between these products are optimized to fit the experimental data. 239 These branching ratios cannot be compared directly to previously reported VBS 240 241 parameters for  $\alpha$ -pinene ozonolysis (e.g., Henry et al. (2012)) since VBS parameters are 242 typically mass-based, while the branching ratios in the model are mole-based. 243 Furthermore, the branching ratios here account for the influence of vapor wall deposition, 244 while typical VBS parameters do not. We assume that these 5 classes of products have molecular weights 168, 184, 192, 200 and 216 g mole<sup>-1</sup> based on the group contribution 245 246 method (Donahue et al., 2011). The first-generation products are assumed not to undergo 247 further reaction with O<sub>3</sub> upon formation.

The aerosol dynamics in the chamber obey the aerosol general dynamic equation(Seinfeld and Pandis, 2006):

250 
$$\left(\frac{\partial n(D_p,t)}{\partial t}\right) = \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{coag}} + \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{cond/evap}} + \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{wall loss}}$$
(5)

251 Coagulation is not considered, since an alternative version of the model including 252 coagulation showed no change in the predicted  $\alpha$ -pinene ozonolysis SOA mass 253 concentrations in simulations with and without coagulation. The change in particle 254 number distribution due to particle wall loss is:

255 
$$\left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{wall loss}} = -\beta_j(D_p)n(D_p,t)$$
(6)

where, as noted in section 2.2,  $\beta_j(D_p)$  is the size-dependent first-order wall loss rate coefficient obtained from fitting seed-only experiments. The rate at which vapor condenses onto a spherical aerosol particle is:

259 
$$J_{i} = 2\pi D_{p} D_{i} (G_{i} - G_{i}^{eq}) F_{FS}$$
(7)

where  $G_i$  is the concentration of gas-phase species *i*,  $G_i^{eq}$  is the saturation concentration of gas-phase species *i*,  $D_i$  is the gas-phase molecular diffusivity (assumed to be  $3 \times 10^{-6}$ 





262  $m^2 s^{-1}$  (McVay et al., 2014)), and  $F_{FS}$  is the Fuchs-Sutugin correction for non-continuum 263 gas-phase diffusion:

264 
$$F_{FS} = \frac{0.75\alpha_p (1 + \text{Kn})}{\text{Kn}^2 + \text{Kn} + 0.283\text{Kn}\alpha_p + 0.75\alpha_p}$$
(8)

where  $\alpha_p$  is the vapor-particle mass accommodation coefficient, and Kn is the Knudsen number, Kn =  $2\lambda_{AB}/D_p$ . The vapor-particle mass accommodation coefficient accounts for any resistance to vapor molecule uptake at the particle surface (e.g. surface accommodation and particle-phase diffusion limitations).  $\lambda_{AB}$  is the mean free path of the gas-phase species, which is:

270 
$$\lambda_{AB} = 3D_i \times \sqrt{\frac{\pi M_i}{8RT}}$$
(9)

where *R* is the ideal gas constant, *T* is the temperature, and  $M_i$  is the molecular weight of diffusing gas-phase molecule *i*. For each particle size bin, Eqs. 7-9 are used to compute the flux of each gas-phase species to and from an aerosol particle, scaled by the particle number concentration in the size bin. The net rate of change for each gas-phase species due to evaporation or condensation is obtained from the total flux summed over all the particle size bins.

277  $G_i^{eq}$  varies for each particle size bin because it depends on the mass concentration 278 of species *i* and the total organic mass concentration in the size bin:

279 
$$G_{i}^{eq} = \frac{A_{i}C_{i}^{*}}{\sum_{k} A_{k} + M_{init}}$$
(10)

where  $A_i$  is the concentration of species *i* in the particle phase,  $C_i^*$  is the saturation concentration of species *i*,  $\sum_k A_k$  is the sum of all the species concentration in the particle phase, and  $M_{init}$  is the mass concentration of any absorbing organic material initially





283 present in the seed aerosol. To avoid numerical errors in Eq. 10 at the first time step,  $M_{init}$ 

284 is set to 0.01  $\mu$ g m<sup>-3</sup>.

285 The oxidation products of  $\alpha$ -pinene ozonolysis are assumed to be subject to 286 vapor-wall deposition, which is simulated using a first-order wall-loss coefficient 287 (McMurry and Grosjean, 1985):

288 
$$k_{wall,on} = \left(\frac{A}{V}\right) \frac{\frac{\alpha_{wall}c}{4}}{1 + \frac{\pi}{2}\left(\frac{\alpha_{wall}c}{4\sqrt{k_e D_i}}\right)}$$
(11)

where A/V is the surface area-to-volume ratio of the chamber (estimated to be 2.5 m<sup>-1</sup>),  $\alpha_{wall}$  is the vapor-wall mass accommodation coefficient, and  $k_e$  is the eddy diffusion coefficient that describes mixing conditions in the chamber. Based on the measured sizedependent particle wall loss rates (method is described in Zhang et al. (2014)),  $k_e$  is estimated to be 0.03 s<sup>-1</sup> for the GTEC chamber. Vapor-wall deposition is assumed to be reversible, and the rate constant of vapor desorption from the chamber walls is:

295 
$$k_{wall,off} = \frac{k_{wall,on}}{K_w C_w} = k_{wall,on} \left(\frac{C_i^* M_w \gamma_w}{C_w M_p \gamma_p}\right)$$
(12)

296 where  $C_w$  is the equivalent organic mass concentration in the wall (designated to treat 297 gas-wall partitioning in terms of gas-particle partitioning theory and not necessarily 298 representative of a physical layer of organic concentration on the wall (Matsunaga and 299 Ziemann, 2010)),  $K_w$  is the gas-wall partitioning coefficient,  $M_w$  is the effective molecular weight of the wall material,  $\gamma_w$  is the activity coefficient of the species in the wall layer, 300  $M_p$  is the average molecular weight of organic species in the particle, and  $\gamma_p$  is the activity 301 coefficient of the species in the particle. For simplicity, we assume that  $M_w = M_p$  and  $\gamma_w =$ 302  $\gamma_p$ .  $C_w$  is set to 10 mg m<sup>-3</sup> based on previous inferences by Matsunaga and Ziemann 303 304 (2010). Sensitivity studies (not shown) show no change in model predictions when 305 varying  $C_w$  above  $C_w = 0.1 \text{ mg m}^{-3}$ .





In the initial version of the model, after all the  $\alpha$ -pinene is consumed, vapor wall 306 307 deposition was assumed to continue to deplete the gas-phase oxidation products and 308 aerosol mass evaporates to maintain gas-particle equilibrium. SOA evaporation was not 309 observed experimentally (i.e., the SOA mass concentration does not decrease 310 significantly over time after peak SOA growth has been achieved in these chamber 311 experiments (Fig. 2)). In order to represent these observations in the model, a first-order, 312 particle-phase reaction is introduced by which aerosol species are converted into nonvolatile absorbing organic mass with a timescale of  $\tau_{olig}$ . This mechanism (which is not 313 314 included in the model used in McVay et al. (2014)) is similar to that used by the 315 sequential equilibrium partitioning model, in which aerosol is converted from an 316 absorbing to non-absorbing, non-volatile phase in order to explain the inhibited diffusion 317 and evaporation observed in  $\alpha$ -pinene ozonolysis SOA (Cappa and Wilson, 2011). 318 Although we assume here that the converted non-volatile aerosol mass still participates in 319 partitioning, either mechanism invokes a particle-phase process to retard SOA 320 evaporation.

321 Model parameters  $\alpha_w$ ,  $\alpha_p$ ,  $\tau_{olig}$  and the branching ratios between the oxidation 322 products are optimized to best-fit the predictions with the experimental observations. Specifically, model predictions are compared to experimental data: SOA mass 323 324 concentration vs. reaction time,  $\alpha$ -pinene concentration vs. reaction time, and O<sub>3</sub> concentration vs. reaction time. Figure S6 compares reaction profiles of measured and 325 326 modeled  $O_3$  and  $\alpha$ -pinene concentrations for the base model case. Sensitivity tests were 327 also performed on each model parameter, shown in Figs, S7-S10. Table 2 summarizes the 328 parameters used. While the optimized parameters provide a good fit to the data, we 329 caution that the parameters are interconnected, and other fits may also be possible. We 330 are confident that our conclusions derived using these parameters are robust.

### **331 4. Results**

Figure 1 shows the size-dependent particle wall deposition coefficients inferred from seed-only deposition experiments. The initial total AS seed surface area concentration in the low AS-seed only and high AS-seed only experiments (which are conducted using 0.015 M AS and 0.05 M AS solutions, respectively) are similar to those





used in the  $\alpha$ -pinene ozonolysis experiments (i.e., ~1000 and ~3000  $\mu$ m<sup>2</sup> cm<sup>-3</sup>, respectively). While there are differences in the particle wall deposition coefficients from the low and high AS-seed only experiments, this difference is likely the result of uncertainties arising from the low particle number concentrations for the larger particles in the low AS-seed only experiment. As shown in Fig. 1, both sets of particle wall deposition coefficients generally fall within the range of those measured in routine monthly AS-seed only experiments conducted in the chamber.

343 The particle wall deposition corrected number concentration data provide a test of 344 the appropriateness of the particle wall deposition correction. This is because the 345 corrected number concentration should level off at a constant value (i.e., the initial 346 particle number concentration), assuming no significant coagulation, when particle wall 347 deposition is properly accounted for since the wall-deposited particle number distribution 348 is added to the suspended particle number distribution during particle wall loss 349 correction. We account for particle wall deposition in nucleation and low AS experiments 350 using deposition coefficients determined from the low AS-seed only experiments, while 351 particle deposition in high AS experiments are accounted for using coefficients 352 determined from the high AS-seed only experiments. Figures S1 and S2 show the particle 353 wall deposition-corrected aerosol number and volume concentrations. Over all 354 experiments, the particle wall deposition-corrected final particle number concentration (i.e., at the end of the reaction) is 9 to 17 % less than the initial particle number 355 concentration for the low AS and high AS experiments (Table S1), respectively, 356 357 indicating that the particle wall deposition-corrected volume concentrations are slightly 358 underestimated. It is currently unclear why the particle wall deposition-corrected final 359 particle number concentrations are somewhat smaller than the initial particle number concentrations, though this could be due to variations in particle wall deposition rates in 360 361 the AS-seed only and  $\alpha$ -pinene ozonolysis experiments. As a sensitivity test, we used the 362 average of the low AS-seed only and high AS-seed only particle wall deposition 363 coefficients to account for particle wall deposition in all the experiments (Figs. S3 and 364 S4). While there is a negligible difference in the particle wall deposition corrected 365 volume concentrations (Figs. S3 and S4 vs. Figs. S1 and S2), a larger spread (1 to 22 %) 366 exists in the difference between the initial and final particle number concentrations when





367 the average particle wall deposition coefficients are used (Table S1). Therefore, all 368 subsequent nucleation and low AS data presented here are particle wall deposition-369 corrected using coefficients determined from the low AS-seed only experiments, and all 370 high AS data are corrected using particle wall deposition coefficients determined from 371 the high AS-seed only experiments.

372 Figure 2 shows the reaction profiles of the  $\alpha$ -pinene ozonolysis experiments. SOA 373 growth typically starts within 10 to 20 min of the start of the reaction. At either initial O<sub>3</sub> 374 concentration, the molar ratio of  $O_3$  reacted to  $\alpha$ -pinene reacted is approximately 1:1 (i.e., 50 ppb  $\alpha$ -pinene reacted with 50 ppb O<sub>3</sub>), which indicates that O<sub>3</sub> reacts only with  $\alpha$ -375 376 pinene and not its oxidation products. As anticipated, the  $\alpha$ -pinene oxidation rates in the 377 100 ppb O<sub>3</sub> experiments are significantly slower than those in the 500 ppb O<sub>3</sub> 378 experiments. Figures 2a-c show that peak SOA levels are typically reached at reaction 379 time ~300 to 350 min in the 100 ppb O<sub>3</sub> experiments, during which  $\ge$  95 % of the injected  $\alpha$ -pinene has reacted. In contrast, all the  $\alpha$ -pinene reacts within 80 to 90 min of 380 381 the start of reaction in the 500 ppb O<sub>3</sub> experiments, and peak SOA levels are achieved at reaction time ~100 min (Figs. 2d-f). These results indicate that the O<sub>3</sub> concentration 382 383 dictates both the rate of  $\alpha$ -pinene oxidation and the time it takes to achieve peak SOA 384 growth.

385 Figure 3 shows the time-dependent growth curves (SOA mass concentration vs. 386  $\alpha$ -pinene reacted (Ng et al., 2006)) for the 100 and 500 ppb O<sub>3</sub> experiments. Only SOA 387 growth data up to SOA peak concentrations are shown. SOA growth essentially stops 388 once all the  $\alpha$ -pinene has reacted. This is expected, as  $\alpha$ -pinene has only one double 389 bond; the first step of  $\alpha$ -pinene ozonolysis is rate-limiting and the first-generation 390 products are condensable (Ng et al., 2006; Chan et al., 2007). The time-dependent SOA 391 growth curves for experiments corresponding to different seed aerosol concentrations 392 overlap for both low and high O<sub>3</sub> concentrations. This indicates that the initial AS seed 393 surface area does not influence the SOA growth rate within the range of AS seed surface 394 area concentration used. It is important to note that while it appears that the SOA growth 395 rate is faster in the 100 ppb O<sub>3</sub> relative to the 500 ppb O<sub>3</sub> experiments based on the time-396 dependent growth curves shown in Fig. 3, this is not the case. Instead, the observed time-





397 dependent growth curves can be explained by the higher concentration of  $\alpha$ -pinene 398 having reacted during the 10 to 20 min delay of SOA formation in the 500 ppb O<sub>3</sub> 399 experiments compared to the 100 ppb O<sub>3</sub> experiments (Fig. 2).

400 Figure 4 shows the time-dependent SOA mass yields as a function of initial total 401 AS seed surface area for the 100 and 500 ppb  $O_3$  experiments. Regardless of the  $O_3$ 402 concentration, the SOA mass yields stay roughly constant despite the increase in AS seed 403 surface area. This indicates that the surface area concentration of AS seed aerosol does 404 not noticeably influence the partitioning of gas-phase  $\alpha$ -pinene ozonolysis products to the 405 particle phase within the range of AS seed surface area concentration used. Higher SOA 406 mass yields are observed in the 500 ppb  $O_3$  experiments, which indicates that the  $\alpha$ -407 pinene oxidation rate controls the absolute amount of SOA formed. It is important to note 408 that these conclusions are robust even when the average of the low AS-seed only and 409 high AS-seed only particle wall loss coefficients are used to account for particle wall loss 410 in all the experiments (Fig. S5). The enhancement of SOA mass yields at higher O<sub>3</sub> 411 concentrations and the lack of a SOA mass yield dependence on AS seed surface area 412 (within the range of AS seed surface area concentration used in this study) will be 413 discussed further in Section 5.

The  $\alpha$ -pinene ozonolysis SOA mass yields obtained in this study are compared to those reported in previous studies in Fig. 5. Table S2 lists the experimental conditions employed in these studies. To facilitate comparison between the different studies, all the SOA mass yield and concentration data (including this study) are adjusted to an organic density of 1.0 g cm<sup>-3</sup>. As shown in Fig. 5, the SOA mass yields obtained at peak SOA growth in this study are generally consistent with those of previous studies where the chamber was operated in batch mode (that in this study).

The competition between the condensation of SOA-forming vapor to aerosol particles vs. to chamber walls is investigated using the coupled vapor-particle dynamics model described in Section 3. As noted earlier, optimal model values for  $\alpha_p$ ,  $\alpha_w$ ,  $\tau_{olig}$  and the branching ratios between the oxidation products were determined for the 100 and 500 ppb O<sub>3</sub> experiments by comparing the observed and best-fit time-dependent SOA,  $\alpha$ pinene, and O<sub>3</sub> concentrations profiles (Figs. S6-S10). Sensitivity tests were performed





427 for each parameter to establish that the set of optimal model values provide the best 428 overall agreement with time-dependent SOA formation profiles observed for all experiments (Figs. S7-S10). Predictions from the coupled vapor-particle dynamics model 429 430 show that the optimal parameters are:  $\alpha_w = 10^{-6}$ ,  $\alpha_p = 1$ ,  $\tau_{olig} = 4$  h, branching ratios = 0.57, 0.35, 0.04, 0.015 and 0.025 for oxidation products with vapor pressures  $>10^3$ ,  $10^2$ , 10, 1 431 and 0.1 µg m<sup>-3</sup>, respectively. The best-fit  $\alpha_w = 10^{-6}$  (Fig. S7) corresponds to a first-order 432 vapor-wall deposition rate constant  $(k_{wall,on})$  of  $10^{-4}$  s<sup>-1</sup>. This  $k_{wall,on}$  value is comparable to 433 that reported by Matsunaga and Ziemann (2010) for a 8.2 m<sup>3</sup> chamber. 434

# 435 **5. Discussion**

436  $\alpha$ -pinene ozonolysis has been carried out at two O<sub>3</sub> mixing ratios (100 and 500 437 ppb) under varying AS seed aerosol surface area concentrations (0, ~1000 and ~3000  $\mu$ m<sup>2</sup> 438 cm<sup>-3</sup>).

439 **5.1 Seed aerosol surface area effect** 

440 Figure 3 shows that the time-dependent SOA growth curves for experiments with 441 different seed area concentrations overlap at both O<sub>3</sub> concentrations, which indicates the 442 AS seed surface area does not affect the rate of SOA growth within the range of AS seed 443 surface area concentration used in this study. This observation differs from findings by 444 Pathak et al. (2007b) for the  $O_3+\alpha$ -pinene system, who showed that even though the final SOA mass yields measured in the reaction of 7.3 ppb  $\alpha$ -pinene with 1500 ppb O<sub>3</sub> were 445 446 similar in their seeded and unseeded experiments, SOA growth was considerably slower 447 in unseeded experiments compared to seeded experiments. The authors suggested that the 448 slow SOA formation rate in their unseeded experiment was the result of SOA formation 449 being limited by the mass transfer of semi-volatile oxidation products to newly formed 450 particles (via nucleation) during the early stages of the experiment. These newly formed 451 particles have a significantly smaller aerosol surface area for gas-particle partitioning as 452 compared to that of seed aerosol particles in the seeded experiments. Consequently, the 453 semi-volatile oxidation products accumulated in the gas phase during the early stages of 454 the unseeded experiments, resulting in slower SOA growth compared to the seeded 455 experiments. The observation that the presence of seed aerosol does not influence the





456 SOA growth rate in the present study may be explained by the relatively high 457 concentrations of a-pinene reacted and SOA mass loadings obtained. Previous studies 458 have shown that the delay between the onset of VOC oxidation and SOA formation in 459 unseeded experiments is most pronounced at low aerosol loadings (Kroll et al., 2007). 460 We note that the concentrations of  $\alpha$ -pinene reacted and SOA mass loadings obtained in 461 this study are significantly larger than those reported by Pathak et al. (2007b). Therefore, 462 it is possible that due to the relatively large concentrations of  $\alpha$ -pinene reacted in this 463 study, substantial concentrations of gas-phase oxidation products are generated, which 464 results in rapid partitioning into the particle phase even in the absence of seed aerosol. 465 This is evident from the large increase in the particle number concentration during the early stages of the unseeded 100 and 500 ppb O<sub>3</sub> experiments, where the particle number 466 concentration increased to ~8000 and ~10000 particles/cm<sup>3</sup> during the first 45 min of the 467 468 100 and 500 ppb O<sub>3</sub> experiments, respectively (Fig. S1a and S2a). Thus, the SOA growth 469 rates are not controlled by the presence of AS seed in this study.

Figure 4 shows that for both initial  $O_3$  mixing ratios used, the time-dependent SOA mass yield is similar at any given AS seed surface area (see also Table 1). The absence of a SOA growth dependence on the AS seed surface area is similar to observations reported by McVay et al. (2016) for the  $\alpha$ -pinene photooxidation (OHdriven chemistry) system, but differ from those reported by Zhang et al. (2014) for the toluene photooxidation system in which the SOA mass yield increased with the surface area concentration of seed aerosol.

477 The best-fit  $\alpha_p = 1$  (Fig. S8) suggests the absence of significant limitations to 478 vapor-particle mass transfer in the present  $\alpha$ -pinene ozonolysis study, and that SOA 479 formation is governed by quasi-equilibrium growth (McVay et al., 2014), which occurs 480 when SOA-forming vapors are produced at a rate that is significantly slower than that 481 required to establish gas-particle equilibrium (Shiraiwa and Seinfeld, 2012; Zhang et al., 482 2012). Moreover, the characteristic timescale to establish gas-particle equilibrium is less 483 than those for reaction and vapor-wall deposition. When the vapor and particle phases 484 maintain equilibrium, gas-particle equilibrium is controlled by the amount of organic 485 matter in the VOC system. As a result, the rate of condensation of SOA-forming vapors





is independent of the seed aerosol surface area (McVay et al., 2014). The best-fit  $\alpha_p = 1$  is 486 487 within the approximate range of  $\alpha_p$  coefficients determined from  $\alpha$ -pinene ozonolysis 488 SOA thermodenuder studies ( $\alpha_p = 0.1$ ) (Saleh et al., 2013) and  $\alpha$ -pinene photooxidation 489 chamber studies ( $\alpha_p = 0.1$  or 1) (McVay et al., 2016). Notably, this result differs markedly 490 from that for toluene photooxidation (Zhang et al., 2014), where  $\alpha_p$  was determined to be 491 0.001, and for which, since the SOA mass yield was strongly dependent on the seed 492 aerosol surface area, the condensation of SOA-forming vapors onto seed aerosol particles 493 was kinetically limited (McVay et al., 2014). Kinetically-limited SOA growth occurs 494 when the timescale for gas-particle equilibrium is competitive with or exceeds the 495 timescale for reaction and vapor wall deposition, and may reflect imperfect 496 accommodation of gas-phase organics to the particle phase. The markedly different 497 behavior of the  $\alpha$ -pinene and toluene SOA systems could be due to differences in SOA 498 volatility and aerosol physical phase state (McVay et al., 2016).

# 499 **5.2 Oxidation rate effect**

500 At higher O<sub>3</sub> concentrations, the  $\alpha$ -pinene oxidation rate increases, leading to 501 higher SOA mass yields (the "oxidation rate effect"). This behavior was previously 502 observed by Ng et al. (2007) for the *m*-xylene photooxidation system, for which the 503 oxidation rate effect was attributed to the loss of semi-volatile condensable products to 504 chamber walls in competition with condensation onto seed particles to form SOA.

505 SOA formation from  $\alpha$ -pinene ozonolysis is presumed to be driven by a range of 506 semi- and low-volatility first-generation products arising from reaction of O<sub>3</sub> with the 507 single C=C double bond (Ng et al., 2006). These products are subject to two competing 508 routes: condensation to particles to form SOA or deposition on the chamber walls. Each 509 process can be represented in terms of a first-order rate constant:  $k_{wall,on}$  and  $k_{particle,on}$  (s<sup>-1</sup>). The rate of vapor-wall deposition of condensable species A is then  $k_{wall,on} \times [A]$  (molec 510 cm<sup>-3</sup> s<sup>-1</sup>) and the rate of condensation onto particles is  $k_{particle,on} \times [A]$  (molec cm<sup>-3</sup> s<sup>-1</sup>). 511 Increasing the rate of reaction increases the concentration of [A], but the relative rates of 512 513 vapor-wall deposition and condensation onto particles will remain the same. In general, however, both vapor-wall deposition and vapor-particle condensation are reversible 514





- 515 processes (McVay et al., 2014; Zhang et al., 2014). The first-order rate constant for
- 516 evaporation from the wall can be represented as (Matsunaga and Ziemann, 2010):

517 
$$k_{wall,off} = k_{wall,on} \left(\frac{C_i^*}{C_w}\right)$$
(13)

518 where  $C_i^*$  is the saturation concentration and  $C_w$  is the assumed equivalent wall organic

519 concentration. The rate of evaporation from particles is:

520 
$$k_{particle,off} = k_{particle,on} \left(\frac{C_i^*}{C_{aer}}\right)$$
(14)

521 where  $C_{aer}$  is the organic aerosol concentration ( $C_{aer} = \sum A_k + M_{init}$ ).

522 The difference between  $C_{aer}$  and  $C_w$  is the key to explaining the oxidation rate effect. At the beginning of the experiment, Caer is very small because the inorganic seeds 523 are essentially non-absorbing. Therefore,  $k_{particle,off}$  is large, and the net SOA growth is 524 525 small. In contrast,  $C_w$  is considered to be substantial (on the order of 10 mg m<sup>-3</sup>) and to be 526 essentially constant throughout the experiment (Matsunaga and Ziemann, 2010; McVay 527 et al., 2014; Zhang et al., 2014). Model predictions are insensitive to the value of  $C_w$ 528 since, in any event,  $C_w$  is significantly larger than  $C_{aer}$  (Zhang et al., 2014). Therefore, 529  $k_{wall,off}$  is small at the beginning of the experiment and the net vapor wall loss rate is fast. 530 As Caer increases, the net SOA condensation rate increases relative to the net vapor wall 531 loss rate. When the reaction rate increases corresponding to higher  $O_3$  concentrations,  $C_{aer}$ 532 grows more quickly because more condensable species are available to form SOA, and 533 the net condensation rate increases more rapidly. Therefore, the observed oxidation rate 534 effect is due to vapor wall deposition, and arises because vapor-particle and vapor-wall 535 condensation are essentially reversible processes. This explanation is consistent with 536 simulations varying the O<sub>3</sub> concentration in which all species are non-volatile (i.e., do not 537 evaporate from the particles or the wall). In this case, no oxidation rate effect is observed as the O<sub>3</sub> concentration increases. The growth curves for different O<sub>3</sub> concentrations 538 539 overlap, and the same yield is obtained regardless of O<sub>3</sub> concentration (Fig. S11).





540 Sensitivity tests were performed to determine the point at which SOA formation is 541 no longer influenced by the  $O_3$  concentration. In these simulations, the initial  $\alpha$ -pinene 542 concentration is fixed at 48 ppb, while the O<sub>3</sub> concentration is varied from 75 to 1000 543 ppb. The rate of O<sub>3</sub> injection is assumed to remain constant as the O<sub>3</sub> concentration is 544 increased to mimic the experimental protocol (i.e.,  $O_3$  injection time is increased to 545 achieve higher O<sub>3</sub> concentrations). The O<sub>3</sub> injection rate used in these simulations is fixed at 500/54.25 ppb min<sup>-1</sup>, which is the same as that used to analyze results from the 500 546 547 ppb O<sub>3</sub> experiments. Model predictions in Fig. S12 show that the maximum SOA mass 548 concentration increases with increasing O<sub>3</sub> concentration up to approximately 500 ppb 549 O<sub>3</sub>. Beyond this O<sub>3</sub> concentration, the SOA growth curves overlap and the maximum 550 SOA mass concentration does not increase even when more  $O_3$  is added. This plateau 551 arises due to the lengthening time required to inject increasing amounts of  $O_3$ . More than 552 1 h is required to inject > 500 ppb of  $O_3$ , and by this time, virtually all of the  $\alpha$ -pinene has 553 reacted. Increasing the O<sub>3</sub> concentration after all of the α-pinene has reacted does not lead 554 to any changes in the SOA mass concentration. However, if a faster injection rate of  $O_3$  is used, the oxidation rate effect will persist to higher  $O_3$  concentrations (i.e., > 500 ppb  $O_3$ ) 555 556 (Fig S13). With a faster injection rate, 500 ppb  $O_3$  is injected before all of the  $\alpha$ -pinene 557 has reacted. Continuing to inject O<sub>3</sub> to a higher concentration (i.e., 750 ppb) will cause  $\alpha$ -558 pinene to decay faster and SOA to grow faster than when the O<sub>3</sub> injection stops at 500 ppb. The oxidation rate effect is then apparent at higher O3 concentrations. If, instead of 559 560 using an injection rate of O<sub>3</sub>, simulations are run using fixed initial O<sub>3</sub> (not possible 561 experimentally), the rate effect persists to even higher  $O_3$  concentrations. The relative 562 increase in yield with increasing O<sub>3</sub> concentrations slows at very high O<sub>3</sub> concentrations 563 because the rate of reaction becomes substantially faster than the vapor wall deposition 564 rate, and there is less marginal effect to increasing the reaction rate.

# 565 5.3 Interplay of the seed aerosol surface area effect and the oxidation rate effect

566 In this study, we observe an oxidation rate effect but not a seed aerosol surface 567 area effect. In Zhang et al. (2014), a seed aerosol surface area effect was observed, but 568 the variation of the oxidation rate was not studied. A key aspect of vapor wall deposition 569 is the potential interplay between the seed aerosol surface area effect and the oxidation





rate effect. To examine this interplay in the  $\alpha$ -pinene ozonolysis system, simulations were 570 571 carried out by varying the seed aerosol surface area and the O<sub>3</sub> concentration 572 simultaneously, while using the branching ratios, oligomerization rate, and vapor wall deposition rate parameters obtained in the present study. The initial  $\alpha$ -pinene 573 574 concentration was set to 50 ppb, and a fixed  $O_3$  concentration was used in place of a 575 linear injection.  $\alpha_p$  was varied at 0.001, 0.01, and 1 in these simulations. Figure 6 shows 576 the SOA mass yield at peak SOA growth as a function of both the seed aerosol surface 577 area and O<sub>3</sub> concentration for  $\alpha_p = 1, 0.01$ , and 0.001. For  $\alpha_p = 1$ , the oxidation rate 578 dominates: SOA mass yield increases significantly as O<sub>3</sub> concentration increases while 579 the seed aerosol surface area has a negligible effect. For  $\alpha_p = 0.01$ , both effects can be 580 observed in different regions: at low O<sub>3</sub> concentrations and high seed aerosol surface 581 areas, the oxidation rate effect dominates; at low seed aerosol surface areas and high  $O_3$ 582 concentrations, the seed surface area dominates. At low seed aerosol surface areas and 583 low O<sub>3</sub> concentrations, both effects are present. For  $\alpha_p = 0.001$ , the seed aerosol surface 584 area effect dominates except at very high seed aerosol surface areas. These observations 585 show that the presence of an oxidation rate effect and/or seed aerosol surface area effect depends on a complex interplay of factors, such as  $\alpha_p$ , the rate of hydrocarbon oxidation, 586 587 and the amount of seed surface area present.

### 588 6. Implications

589 In this study, we systematically examine the roles of gas-particle partitioning and 590 VOC oxidation rate in the presence of vapor-wall deposition in  $\alpha$ -pinene ozonolysis. We 591 show that despite the presence of vapor-wall deposition, SOA mass yields at peak SOA 592 growth remain approximately constant regardless the seed aerosol surface area (within 593 the range of AS seed surface area concentration used in this study). This observation is 594 consistent with SOA formation in the  $\alpha$ -pinene ozonolysis system being governed by 595 quasi-equilibrium growth, for which there are no substantial limitations to vapor-particle 596 mass transfer. This result was demonstrated in a previous modeling study which showed 597 that increasing the seed-to-chamber surface area ratio will lead to increased SOA growth 598 only in cases in which the condensation of SOA-forming vapors onto seed aerosol





599 particles is kinetically limited as a result of imperfect accommodation of gas-phase 600 organics to the particle phase (McVay et al., 2014).

An important implication of this study is that diverting vapor-wall deposition in chamber studies via the addition of ever-increasing quantities of seed aerosol particles is not effective in VOC systems for which SOA formation is governed by quasi-equilibrium growth. This study also underscores the importance of accounting for particle wall deposition appropriately in chamber studies, to avoid erroneous conclusions regarding the role of gas-particle partitioning (quasi-equilibrium vs. kinetically-limited SOA growth) in influencing vapor wall loss in the VOC system.

608 We note that the present study shows that the SOA mass yield is independent of seed aerosol surface area concentration for values ranging from 0 to  $\sim 3000 \ \mu m^2 \ cm^{-3}$ . 609 This corresponds to a seed-to-chamber surface area ratio of 0 to  $\sim 1 \times 10^{-3}$ , which is 610 611 substantially smaller than the range used by Zhang et al. (2014) to study the influence of 612 vapor-wall deposition on toluene photooxidation SOA formation in the Caltech chamber (i.e., 0 to  $\sim 5 \times 10^{-3}$ ). It is possible that a SOA mass yield dependence on the seed surface 613 area may have become more apparent had a larger range of seed aerosol surface area (i.e., 614  $> 3000 \ \mu\text{m}^2 \ \text{cm}^{-3}$ ), and hence a larger range of seed-to-chamber surface area ratio, been 615 616 used here. One consideration is that coagulation may become increasingly important, and 617 will need to be accounted for, when higher seed aerosol number concentrations (relative to those used in this study) are used (Seinfeld and Pandis, 2006; Pierce et al., 2008). A 618 619 detailed analysis of the effect of seed aerosol surface area concentrations  $> 3000 \ \mu m^2 \ cm^2$ <sup>3</sup> on  $\alpha$ -pinene ozonolysis SOA mass yields will be the subject of forthcoming work. 620

621 Higher SOA mass yields at peak SOA growth are observed in the present study 622 when  $O_3$  is increased from 100 to 500 ppb. This is because  $\alpha$ -pinene is oxidized more 623 quickly, which leads to gas-phase oxidation products being formed more rapidly, and 624 consequently partitioning more quickly onto AS seed aerosol particles before they are 625 lost to the chamber walls. Therefore, the oxidation rate effect (i.e., higher SOA mass 626 vields as a result of faster hydrocarbon oxidation rates) is a consequence of vapor-wall 627 deposition. An important implication of this study is that SOA mass yields can be 628 affected by vapor-wall deposition in VOC systems that are not characterized by slow





629 mass accommodation of gas-phase organics to the particle phase (Zhang et al., 2014). 630 Thus, this work demonstrates that the effect of vapor-wall deposition on SOA mass yields 631 can be mitigated through the use of excess oxidant concentrations. It should be noted that 632 the  $\alpha$ -pinene ozonolysis SOA mass yields (absolute values) increased by 5 to 9 % when 633  $O_3$  is increased from 100 to 500 ppb (for an initial  $\alpha$ -pinene concentration of ~50 ppb), 634 where SOA formation is governed by quasi-equilibrium growth. In the absence of vapor-635 wall deposition, SOA mass yields are predicted by the model used here to approximately 636 double from those observed experimentally. In contrast, Zhang et al. (2014) showed that 637 the presence of vapor-wall deposition led to underestimation of SOA formation by factors 638 as much as four in the toluene photooxidation system, where the condensation of SOA-639 forming vapors onto seed aerosol is kinetically limited. Taken together, these results 640 indicate that the magnitude by which vapor-wall deposition affects SOA mass yields 641 depends on the extent to which the VOC system is governed by kinetically-limited SOA 642 condensational growth.

643 Given these observations of how gas-particle partitioning can influence the 644 magnitude by which vapor-wall deposition affects SOA mass yields, an overriding 645 question is: what controls the gas-particle partitioning behavior of SOA formed in 646 different VOC systems?  $\alpha_p$  describes the overall mass transfer of vapor molecules into the 647 particle phase (McVay et al., 2014; Zhang et al., 2014). Thus,  $\alpha_p$  affects the vapor-648 particle equilibrium timescale, which, depending on the extent to which it is competitive 649 with the timescales for reaction and vapor-wall deposition, determines whether SOA 650 formation is governed by kinetically-limited or quasi-equilibrium growth. Markedly different  $\alpha_p$  values could arise from the physical phase state of the SOA formed. As 651 652 discussed by McVay et al. (2014, 2016), if the SOA formed exists in a semi-solid state 653 (Vaden et al., 2010; Virtanen et al., 2010; Cappa and Wilson, 2011; Vaden et al., 2011; 654 Virtanen et al., 2011; Kuwata and Martin, 2012; Perraud et al., 2012; Saukko et al., 2012; 655 Abramson et al., 2013; Renbaum-Wolff et al., 2013), a low value of  $\alpha_p$  might be expected 656 owing to retarded surface accommodation and particle-phase diffusion (Zaveri et al., 657 2014). Quantification of  $\alpha_p$  is challenging experimentally, and reported  $\alpha_p$  values for the 658 same system can vary by several orders of magnitude (Grieshop et al., 2007; Stanier et 659 al., 2007; Vaden et al., 2011; Miles et al., 2012; Saleh et al., 2013). Therefore,  $\alpha_p$  of SOA





660 formed in different VOC systems need to be better constrained through a combination of

661 experimental and modeling efforts.

662 The SOA mass yield from the ozonolysis of monoterpenes in the GEOS-CHEM chemical transport model (19 % at 10  $\mu$ g m<sup>-3</sup>) is currently based on that measured in  $\alpha$ -663 pinene ozonolysis studies by Shilling et al. (2008) (Pve et al., 2010). Shilling et al. (2008) 664 665 measured these SOA mass yields in a teflon chamber operated in continuous-flow mode, as opposed to batch mode, which is how experiments in the present study and most of 666 those shown in Fig. 5 and Table S2 were conducted. While it is not possible to directly 667 compare our results with those of Shilling et al. (2008) due to differences in SOA mass 668 669 concentrations, the SOA mass concentrations and yields measured in the current study 670 are generally consistent with those of previous batch chamber studies. The SOA mass yields at  $\sim 10 \ \mu g \ m^{-3}$  SOA mass concentration measured by Shilling et al. (2008) are 671 generally higher than those measured in chambers operated in batch mode (Griffin et al., 672 673 1999; Cocker et al., 2001b; Presto et al., 2005; Presto and Donahue, 2006; Pathak et al., 674 2007b) (Fig. 5). One possible explanation for the higher SOA mass yields in the 675 continuous-flow, steady state, mode is that the SOA-forming vapors are in equilibrium 676 with the organic mass present on the chamber walls and seed aerosol, hence minimizing 677 the irreversible loss of SOA-forming vapors to the chamber walls (Shilling et al., 2008). 678 However, the extent to which SOA mass yields obtained in a continuous-flow reactor are 679 influenced by vapor wall loss is unclear. Using a continuous-flow reactor, Ehn et al. (2014) observed  $\alpha$ -pinene ozonolysis SOA mass yields to increase with increasing seed 680 aerosol surface area but required  $\alpha_n = 1$  to fit the observed SOA growth. The observed 681 vapor-wall deposition rate constant in their continuous-flow reactor  $(0.011 \text{ s}^{-1})$  is two 682 orders of magnitude larger than that of the GTEC chamber  $(10^{-4} \text{ s}^{-1})$ . The estimated 683 684 timescales for gas-particle and gas-wall partitioning are also approximately equal in their 685 continuous-flow reactor. This indicates that SOA condensational growth is kinetically 686 limited in their continuous-flow reactor even at  $\alpha_p = 1$  (Ehn et al., 2014; McVay et al., 687 2014), which suggests that SOA mass yields measured in their continuous-flow reactor 688 may be significantly affected by vapor-wall deposition.





689 Previous studies on SOA formation from the OH and NO<sub>3</sub> oxidation of biogenic 690 VOCs have similarly reported higher SOA mass yields in the presence of higher oxidant 691 concentrations. For example, in the NO<sub>3</sub> oxidation of  $\beta$ -pinene, Boyd et al. (2015) 692 reported SOA mass yields 10 to 30 % higher than those previously reported by Fry et al. 693 (2009, 2014). In addition to differences in the experimental conditions of the two studies, 694 Boyd et al. (2015) hypothesized that the higher SOA mass yields could also be a result of 695 the higher NO<sub>3</sub> concentrations used in their study (which led to faster  $\beta$ -pinene oxidation 696 rates) compared to those used by Fry et al. (2009, 2014). The oxidation rate effect was 697 also observed in the *m*-xylene photooxidation system, where Ng et al. (2007) showed that 698 the SOA mass yields were dependent on the *m*-xylene oxidation rate, with higher OH 699 concentrations (and hence faster oxidation rates) resulting in higher SOA mass yields 700 Together, these studies show that faster hydrocarbon oxidation rates can alleviate the 701 effects of vapor-wall deposition on SOA mass yields in different VOC systems.

702 This gives rise to the question: should chamber SOA experiments on different 703 VOC systems be performed under as rapid oxidation conditions as possible (i.e., large 704 oxidant concentrations) to reduce the effects of vapor-wall deposition? A recent study by 705 McVay et al. (2016) reported similar SOA growth under low and high OH levels for  $\alpha$ -706 pinene photooxidation. The authors hypothesized that the autoxidation mechanism likely 707 becomes a more important pathway at low OH levels (Crounse et al., 2013), and thus 708 contributes substantially to SOA growth. Therefore, it is possible that certain reaction 709 pathways and mechanisms (which are important in the atmosphere) are biased when 710 unusually high levels of oxidants are used in chamber experiments (e.g. autoxidation). 711 Thus, this underscores the need to design chamber experiments that simultaneously 712 mitigate the magnitude of vapor-wall deposition while ensuring that reaction conditions, 713 and consequently reaction pathways and oxidation products, are atmospherically relevant.

# 714 Acknowledgements

This research was funded by NSF Grants 1455588 and AGS-1523500, and US Environmental Protection Agency STAR grant (Early Career) RD-83540301. This publication's contents are solely the responsibility of the grantee and do not necessarily represent the official views of the US EPA. Further, US EPA does not endorse the





- 719 purchase of any commercial products or services mentioned in the publication. R.C.
- 720 McVay was supported by a National Science Foundation Graduate Research Fellowship
- 721 under Grant No. DGE-1144469.

# 722 References

- Abramson, E., Imre, D., Beranek, J., Wilson, J., and Zelenyuk, A.: Experimental determination of chemical diffusion within secondary organic aerosol particles, Phys.
- 725 Chem. Chem. Phys., 15, 2983-2991, 10.1039/c2cp44013j, 2013.
- Bahreini, R., Keywood, M. D., Ng, N. L., Varutbangkul, V., Gao, S., Flagan, R. C.,
  Seinfeld, J. H., Worsnop, D. R., and Jimenez, J. L.: Measurements of Secondary Organic
  Aerosol from Oxidation of Cycloalkenes, Terpenes, and m-Xylene Using an Aerodyne
  Aerosol Mass Spectrometer, Environmental Science & Technology, 39, 5674-5688,
  10.1021/es048061a, 2005.
- 731 Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and
- 732 Ng, N. L.: Secondary organic aerosol formation from the  $\beta$ -pinene+NO3 system: effect of
- humidity and peroxy radical fate, Atmos. Chem. Phys., 15, 7497-7522, 10.5194/acp-15-
- 734 7497-2015, 2015.
- 735 Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P.,
- Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue,
- 737 N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic
- 738 compounds using aerosol mass spectrometry: characterization, improved calibration, and
- 739 implications, Atmos. Chem. Phys., 15, 253-272, 10.5194/acp-15-253-2015, 2015.
- 740 Capouet, M., Müller, J. F., Ceulemans, K., Compernolle, S., Vereecken, L., and Peeters,
- 741 J.: Modeling aerosol formation in alpha-pinene photo-oxidation experiments, Journal of
- 742 Geophysical Research: Atmospheres, 113, n/a-n/a, 10.1029/2007JD008995, 2008.
- Cappa, C. D., and Wilson, K. R.: Evolution of organic aerosol mass spectra upon heating:
  implications for OA phase and partitioning behavior, Atmospheric Chemistry and
  Physics, 11, 1895-1911, 10.5194/acp-11-1895-2011, 2011.





- 746 Cappa, C. D., Jathar, S. H., Kleeman, M. J., Docherty, K. S., Jimenez, J. L., Seinfeld, J.
- 747 H., and Wexler, A. S.: Simulating secondary organic aerosol in a regional air quality
- 748 model using the statistical oxidation model Part 2: Assessing the influence of vapor
- 749 wall losses, Atmos. Chem. Phys., 16, 3041-3059, 10.5194/acp-16-3041-2016, 2016.
- 750 Carlton, A. G., Bhave, P. V., Napelenok, S. L., Edney, E. D., Sarwar, G., Pinder, R. W.,
- 751 Pouliot, G. A., and Houyoux, M.: Model Representation of Secondary Organic Aerosol in
- 752 CMAQv4.7, Environmental Science & Technology, 44, 8553-8560, 10.1021/es100636q,
- 753 2010.
- Chan, A. W. H., Kroll, J. H., Ng, N. L., and Seinfeld, J. H.: Kinetic modeling of
  secondary organic aerosol formation: effects of particle- and gas-phase reactions of
  semivolatile products, Atmospheric Chemistry and Physics, 7, 4135-4147, 2007.
- Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: State-of-the-art chamber facility for
  studying atmospheric aerosol chemistry, Environmental Science & Technology, 35,
  2594-2601, 10.1021/es0019169, 2001a.
- 760 Cocker, D. R., Clegg, S. L., Flagan, R. C., and Seinfeld, J. H.: The effect of water on gas-
- particle partitioning of secondary organic aerosol. Part I: alpha-pinene/ozone system,
  Atmospheric Environment, 35, 6049-6072, 10.1016/s1352-2310(01)00404-6, 2001b.
- Crounse, J. D., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., and Wennberg, P. O.:
  Autoxidation of Organic Compounds in the Atmosphere, J. Phys. Chem. Lett., 4, 35133520, 10.1021/jz4019207, 2013.
- 766 Crump, J. G., and Seinfeld, J. H.: Turbulent Deposition and Gravitational Sedimentation
  767 of an Aerosol in a Vessel of Arbitrary Shape, Journal of Aerosol Science, 12, 405-415,
  768 10.1016/0021-8502(81)90036-7, 1981.
- DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle
  morphology and density characterization by combined mobility and aerodynamic
  diameter measurements. Part 1: Theory, Aerosol Sci. Technol., 38, 1185-1205,
  10.1080/027868290903907, 2004.





- 773 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
- 774 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.:
- 775 Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, Analytical
- 776 Chemistry, 78, 8281-8289, 10.1021/ac061249n, 2006.
- 777 Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning,
- 778 dilution, and chemical aging of semivolatile organics, Environmental Science &
- 779 Technology, 40, 2635-2643, 10.1021/es052297c, 2006.
- 780 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional
- 781 volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmospheric Chemistry
- 782 and Physics, 11, 3303-3318, 10.5194/acp-11-3303-2011, 2011.
- 783 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M.,
- 784 Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen,
- 785 M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T.,
- 786 Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso,
- 787 M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R.,
- 788 Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol,
- 789 Nature, 506, 476-479, 10.1038/nature13032, 2014.
- Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs,
  H., Dube, W., Mensah, A., dal Maso, M., Tillmann, R., Dorn, H. P., Brauers, T., and
  Cohen, R. C.: Organic nitrate and secondary organic aerosol yield from NO3 oxidation of
- beta-pinene evaluated using a gas-phase kinetics/aerosol partitioning model, Atmospheric
- 794 Chemistry and Physics, 9, 1431-1449, 2009.
- 795 Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkler, P. M., Lawler, M. J., Brown, S. S., Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic 796 797 Aerosol Formation and Organic Nitrate Yield from NO3 Oxidation of Biogenic 798 Environmental Hydrocarbons, Science & Technology, 48. 11944-11953, 799 10.1021/es502204x, 2014.





- 800 Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J. W.,
- 801 Yoo, K. Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle
- 802 phase acidity and oligomer formation in secondary organic aerosol, Environmental
- 803 Science & Technology, 38, 6582-6589, 10.1021/es049125k, 2004.
- 804 Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Is the gas-particle partitioning in
- alpha-pinene secondary organic aerosol reversible?, Geophys. Res. Lett., 34, n/a-n/a,
  10.1029/2007GL029987, 2007.
- Griffin, R. J., Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation
  from the oxidation of biogenic hydrocarbons, J. Geophys. Res.-Atmos., 104, 3555-3567,
  10.1029/1998jd100049, 1999.
- 810 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L.
- 811 K., and Wang, X.: The Model of Emissions of Gases and Aerosols from Nature version
- 812 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions,
- 813 Geoscientific Model Development, 5, 1471-1492, 10.5194/gmd-5-1471-2012, 2012.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
  Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann,
  H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr,
  A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld,
  J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact
  of secondary organic aerosol: current and emerging issues, Atmospheric Chemistry and
  Physics, 9, 5155-5236, 2009.
- Henry, K. M., Lohaus, T., and Donahue, N. M.: Organic Aerosol Yields from alphaPinene Oxidation: Bridging the Gap between First-Generation Yields and Aging
  Chemistry, Environmental Science & Technology, 46, 12347-12354, 10.1021/es302060y,
  2012.
- Hildebrandt, L., Donahue, N. M., and Pandis, S. N.: High formation of secondary organic
  aerosol from the photo-oxidation of toluene, Atmospheric Chemistry and Physics, 9,
  2973-2986, 2009.





- Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and
  Seinfeld, J. H.: Formation of organic aerosols from the oxidation of biogenic
  hydrocarbons, Journal of Atmospheric Chemistry, 26, 189-222,
- 831 10.1023/a:1005734301837, 1997.
- 832 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C.,
- 833 Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P.,
- 834 Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L.,
- 835 Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global
- climate modelling: a review, Atmospheric Chemistry and Physics, 5, 1053-1123, 2005.
- Keywood, M. D., Varutbangkul, V., Bahreini, R., Flagan, R. C., and Seinfeld, J. H.:
  Secondary organic aerosol formation from the ozonolysis of cycloalkenes and related
  compounds, Environmental Science & Technology, 38, 4157-4164, 10.1021/es0353630,
  2004.
- - - -
- 841 Kokkola, H., Yli-Pirila, P., Vesterinen, M., Korhonen, H., Keskinen, H., Romakkaniemi,
- 842 S., Hao, L., Kortelainen, A., Joutsensaari, J., Worsnop, D. R., Virtanen, A., and Lehtinen,
- 843 K. E. J.: The role of low volatile organics on secondary organic aerosol formation,
- 844 Atmospheric Chemistry and Physics, 14, 1689-1700, 10.5194/acp-14-1689-2014, 2014.
- Kostenidou, E., Pathak, R. K., and Pandis, S. N.: An algorithm for the calculation of
  secondary organic aerosol density combining AMS and SMPS data, Aerosol Sci.
  Technol., 41, 1002-1010, 10.1080/02786820701666270, 2007.
- Kristensen, K., Cui, T., Zhang, H., Gold, A., Glasius, M., and Surratt, J. D.: Dimers in
  alpha-pinene secondary organic aerosol: effect of hydroxyl radical, ozone, relative
  humidity and aerosol acidity, Atmospheric Chemistry and Physics, 14, 4201-4218,
  10.5194/acp-14-4201-2014, 2014.
- Kroll, J. H., Chan, A. W. H., Ng, N. L., Flagan, R. C., and Seinfeld, J. H.: Reactions of
  semivolatile organics and their effects on secondary organic aerosol formation,
  Environmental Science & Technology, 41, 3545-3550, 10.1021/es062059x, 2007.





- 855 Kuwata, M., and Martin, S. T.: Phase of atmospheric secondary organic material affects 856 its reactivity, Proc. Natl. Acad. Sci. U. S. A., 109, 17354-17359, 857 10.1073/pnas.1209071109, 2012.
- 858 La, Y. S., Camredon, M., Ziemann, P. J., Valorso, R., Matsunaga, A., Lannuque, V., Lee-

Taylor, J., Hodzic, A., Madronich, S., and Aumont, B.: Impact of chamber wall loss of
gaseous organic compounds on secondary organic aerosol formation: explicit modeling
of SOA formation from alkane and alkene oxidation, Atmos. Chem. Phys., 16, 1417-

- 862 1431, 10.5194/acp-16-1417-2016, 2016.
- Loza, C. L., Chan, A. W. H., Galloway, M. M., Keutsch, F. N., Flagan, R. C., and
  Seinfeld, J. H.: Characterization of Vapor Wall Loss in Laboratory Chambers,
  Environmental Science & Technology, 44, 5074-5078, 10.1021/es100727v, 2010.
- Loza, C. L., Chhabra, P. S., Yee, L. D., Craven, J. S., Flagan, R. C., and Seinfeld, J. H.:
  Chemical aging of m-xylene secondary organic aerosol: laboratory chamber study,
  Atmospheric Chemistry and Physics, 12, 151-167, 10.5194/acp-12-151-2012, 2012.
- Matsunaga, A., and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a
  Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield
  Measurements, Aerosol Sci. Technol., 44, 881-892, 10.1080/02786826.2010.501044,
  2010.
- McMurry, P. H., and Grosjean, D.: Gas and Aerosol Wall Losses in Teflon Film Smog
  Chambers, Environmental Science & Technology, 19, 1176-1182, 10.1021/es00142a006,
  1985.
- McMurry, P. H., and Rader, D. J.: Aerosol Wall Losses in Electrically Charged
  Chambers, Aerosol Sci. Technol., 4, 249-268, 10.1080/02786828508959054, 1985.
- 878 McVay, R. C., Cappa, C. D., and Seinfeld, J. H.: Vapor-Wall Deposition in Chambers:
- Theoretical Considerations, Environmental Science & Technology, 48, 10251-10258,
  10.1021/es502170j, 2014.





- 881 McVay, R. C., Zhang, X., Aumont, B., Valorso, R., Camredon, M., La, Y. S., Wennberg,
- 882 P. O., and Seinfeld, J. H.: SOA formation from the photooxidation of α-pinene:
- systematic exploration of the simulation of chamber data, Atmos. Chem. Phys., 16, 2785-
- 884 2802, 10.5194/acp-16-2785-2016, 2016.
- 885 Miles, R. E. H., Reid, J. P., and Riipinen, I.: Comparison of Approaches for Measuring
- the Mass Accommodation Coefficient for the Condensation of Water and Sensitivities to Uncertainties in Thermophysical Properties, J. Phys. Chem. A, 116, 10810-10825,
- 888 10.1021/jp3083858, 2012.
- 889 Ng, N. L., Kroll, J. H., Keywood, M. D., Bahreini, R., Varutbangkul, V., Flagan, R. C., 890 Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Contribution of first- versus second-891 generation products to secondary organic aerosols formed in the oxidation of biogenic 892 hydrocarbons, Environmental Science & Technology, 40, 2283-2297, 893 10.1021/es052269u, 2006.
- Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.:
  Secondary organic aerosol formation from m-xylene, toluene, and benzene, Atmospheric
  Chemistry and Physics, 7, 3909-3922, 2007.
- 897 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.:
  898 Gas/Particle Partitioning and Secondary Organic Aerosol Yields, Environmental Science
  899 & Technology, 30, 2580-2585, 10.1021/es950943+, 1996.
- Odum, J. R., Jungkamp, T. P. W., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: The
  atmospheric aerosol-forming potential of whole gasoline vapor, Science, 276, 96-99,
  10.1126/science.276.5309.96, 1997a.
- Odum, J. R., Jungkamp, T. P. W., Griffin, R. J., Forstner, H. J. L., Flagan, R. C., and
  Seinfeld, J. H.: Aromatics, reformulated gasoline, and atmospheric organic aerosol
  formation, Environmental Science & Technology, 31, 1890-1897, 10.1021/es9605351,
  1997b.





- 907 Pathak, R. K., Presto, A. A., Lane, T. E., Stanier, C. O., Donahue, N. M., and Pandis, S.
- 908 N.: Ozonolysis of alpha-pinene: parameterization of secondary organic aerosol mass
- 909 fraction, Atmospheric Chemistry and Physics, 7, 3811-3821, 2007a.
- 910 Pathak, R. K., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of alpha-
- 911 pinene at atmospherically relevant concentrations: Temperature dependence of aerosol
- 912 mass fractions (yields), J. Geophys. Res.-Atmos., 112, 8, 10.1029/2006jd007436, 2007b.
- 913 Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L.,
- 214 Zelenyuk, A., Imre, D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B.
- 915 J.: Nonequilibrium atmospheric secondary organic aerosol formation and growth, Proc.
- 916 Natl. Acad. Sci. U. S. A., 109, 2836-2841, 10.1073/pnas.1119909109, 2012.
- Pierce, J. R., Engelhart, G. J., Hildebrandt, L., Weitkamp, E. A., Pathak, R. K., Donahue,
  N. M., Robinson, A. L., Adams, P. J., and Pandis, S. N.: Constraining particle evolution
  from wall losses, coagulation, and condensation-evaporation in smog-chamber
  experiments: Optimal estimation based on size distribution measurements, Aerosol Sci.
  Technol., 42, 1001-1015, 10.1080/02786820802389251, 2008.
- Presto, A. A., Hartz, K. E. H., and Donahue, N. M.: Secondary organic aerosol
  production from terpene ozonolysis. 2. Effect of NOx concentration, Environmental
  Science & Technology, 39, 7046-7054, 10.1021/es050400s, 2005.
- Presto, A. A., and Donahue, N. M.: Investigation of alpha-pinene plus ozone secondary
  organic aerosol formation at low total aerosol mass, Environmental Science &
  Technology, 40, 3536-3543, 10.1021/es052203z, 2006.
- Pye, H. O. T., Chan, A. W. H., Barkley, M. P., and Seinfeld, J. H.: Global modeling of
  organic aerosol: the importance of reactive nitrogen (NOx and NO3), Atmospheric
  Chemistry and Physics, 10, 11261-11276, 10.5194/acp-10-11261-2010, 2010.
- 931 Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B.
- 932 J., Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of alpha-pinene secondary





- 933 organic material and implications for particle growth and reactivity, Proc. Natl. Acad.
- 934 Sci. U. S. A., 110, 8014-8019, 10.1073/pnas.1219548110, 2013.
- Saleh, R., Donahue, N. M., and Robinson, A. L.: Time Scales for Gas-Particle
  Partitioning Equilibration of Secondary Organic Aerosol Formed from Alpha-Pinene
  Ozonolysis, Environmental Science & Technology, 47, 5588-5594, 10.1021/es400078d,
  2013.
- Saukko, E., Lambe, A. T., Massoli, P., Koop, T., Wright, J. P., Croasdale, D. R.,
  Pedernera, D. A., Onasch, T. B., Laaksonen, A., Davidovits, P., Worsnop, D. R., and
  Virtanen, A.: Humidity-dependent phase state of SOA particles from biogenic and
  anthropogenic precursors, Atmospheric Chemistry and Physics, 12, 7517-7529,
  10.5194/acp-12-7517-2012, 2012.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the
  development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric
  degradation of non-aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161180, 10.5194/acp-3-161-2003, 2003.
- Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics : from air pollution
  to climate change, 2nd ed., Wiley, Hoboken, N.J., xxviii, 1203 p. pp., 2006.
- Shilling, J. E., Chen, Q., King, S. M., Rosenoern, T., Kroll, J. H., Worsnop, D. R.,
  McKinney, K. A., and Martin, S. T.: Particle mass yield in secondary organic aerosol
  formed by the dark ozonolysis of alpha-pinene, Atmospheric Chemistry and Physics, 8,
- 953 2073-2088, 2008.
- Shilling, J. E., Chen, Q., King, S. M., Rosenoern, T., Kroll, J. H., Worsnop, D. R.,
  DeCarlo, P. F., Aiken, A. C., Sueper, D., Jimenez, J. L., and Martin, S. T.: Loadingdependent elemental composition of alpha-pinene SOA particles, Atmospheric Chemistry
  and Physics, 9, 771-782, 2009.
- Shiraiwa, M., and Seinfeld, J. H.: Equilibration timescale of atmospheric secondary
  organic aerosol partitioning, Geophys. Res. Lett., 39, 6, 10.1029/2012gl054008, 2012.





- 960 Song, C., Zaveri, R. A., Alexander, M. L., Thornton, J. A., Madronich, S., Ortega, J. V.,
- 961 Zelenyuk, A., Yu, X. Y., Laskin, A., and Maughan, D. A.: Effect of hydrophobic primary
- 962 organic aerosols on secondary organic aerosol formation from ozonolysis of alpha-
- 963 pinene, Geophys. Res. Lett., 34, 5, 10.1029/2007gl030720, 2007.
- 964 Stanier, C. O., Pathak, R. K., and Pandis, S. N.: Measurements of the volatility of
- 965 aerosols from alpha-piniene ozonolysis, Environmental Science & Technology, 41, 2756-
- 966 2763, 10.1021/es0519280, 2007.
- 967 Tsigaridis, K., Daskalakis, N., Kanakidou, M., Adams, P. J., Artaxo, P., Bahadur, R.,
- 968 Balkanski, Y., Bauer, S. E., Bellouin, N., Benedetti, A., Bergman, T., Berntsen, T. K.,
- 969 Beukes, J. P., Bian, H., Carslaw, K. S., Chin, M., Curci, G., Diehl, T., Easter, R. C.,
- 970 Ghan, S. J., Gong, S. L., Hodzic, A., Hoyle, C. R., Iversen, T., Jathar, S., Jimenez, J. L.,
- 971 Kaiser, J. W., Kirkevåg, A., Koch, D., Kokkola, H., Lee, Y. H., Lin, G., Liu, X., Luo, G.,
- 972 Ma, X., Mann, G. W., Mihalopoulos, N., Morcrette, J. J., Müller, J. F., Myhre, G.,
- 973 Myriokefalitakis, S., Ng, N. L., O'Donnell, D., Penner, J. E., Pozzoli, L., Pringle, K. J.,
- 974 Russell, L. M., Schulz, M., Sciare, J., Seland, Ø., Shindell, D. T., Sillman, S., Skeie, R.
- 975 B., Spracklen, D., Stavrakou, T., Steenrod, S. D., Takemura, T., Tiitta, P., Tilmes, S.,
- 976 Tost, H., van Noije, T., van Zyl, P. G., von Salzen, K., Yu, F., Wang, Z., Wang, Z.,
- 977 Zaveri, R. A., Zhang, H., Zhang, K., Zhang, Q., and Zhang, X.: The AeroCom evaluation
- 978 and intercomparison of organic aerosol in global models, Atmos. Chem. Phys., 14,
- 979 10845-10895, 10.5194/acp-14-10845-2014, 2014.
- Vaden, T. D., Song, C., Zaveri, R. A., Imre, D., and Zelenyuk, A.: Morphology of mixed
  primary and secondary organic particles and the adsorption of spectator organic gases
  during aerosol formation, Proc. Natl. Acad. Sci. U. S. A., 107, 6658-6663,
  10.1073/pnas.0911206107, 2010.
- Vaden, T. D., Imre, D., Beranek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation
  kinetics and phase of laboratory and ambient secondary organic aerosol, Proc. Natl.
  Acad. Sci. U. S. A., 108, 2190-2195, 10.1073/pnas.1013391108, 2011.





- 987 Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinen, J., Makela,
- 988 J. M., Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen, A.:
- 989 An amorphous solid state of biogenic secondary organic aerosol particles, Nature, 467,
- 990 824-827, 10.1038/nature09455, 2010.
- 991 Virtanen, A., Kannosto, J., Kuuluvainen, H., Arffman, A., Joutsensaari, J., Saukko, E.,
- 992 Hao, L., Yli-Pirila, P., Tiitta, P., Holopainen, J. K., Keskinen, J., Worsnop, D. R., Smith,
- 993 J. N., and Laaksonen, A.: Bounce behavior of freshly nucleated biogenic secondary
- 994 organic aerosol particles, Atmospheric Chemistry and Physics, 11, 8759-8766,
- 995 10.5194/acp-11-8759-2011, 2011.
- Weitkamp, E. A., Sage, A. M., Pierce, J. R., Donahue, N. M., and Robinson, A. L.:
  Organic aerosol formation from photochemical oxidation of diesel exhaust in a smog
  chamber, Environmental Science & Technology, 41, 6969-6975, 10.1021/es070193r,
  2007.
- Yeh, G. K., and Ziemann, P. J.: Alkyl Nitrate Formation from the Reactions of C-8-C-14
  n-Alkanes with OH Radicals in the Presence of NOx: Measured Yields with Essential
  Corrections for Gas-Wall Partitioning, J. Phys. Chem. A, 118, 8147-8157,
  10.1021/jp500631v, 2014.
- Yeh, G. K., and Ziemann, P. J.: Gas-Wall Partitioning of Oxygenated Organic
  Compounds: Measurements, Structure-Activity Relationships, and Correlation with Gas
  Chromatographic Retention Factor, Aerosol Sci. Technol., 49, 726-737,
  10.1080/02786826.2015.1068427, 2015.
- Zhang, X., Pandis, S. N., and Seinfeld, J. H.: Diffusion-Limited Versus QuasiEquilibrium Aerosol Growth, Aerosol Sci. Technol., 46, 874-885,
  1010 10.1080/02786826.2012.679344, 2012.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
  Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of
  secondary organic aerosol, Proc. Natl. Acad. Sci. U. S. A., 111, 5802-5807,
  1014 10.1073/pnas.1404727111, 2014.





- 1015 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C.,
- 1016 and Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, Atmospheric Chemistry
- 1017 and Physics, 15, 4197-4214, 10.5194/acp-15-4197-2015, 2015a.
- 1018 Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and
- 1019 Seinfeld, J. H.: Formation and evolution of molecular products in alpha-pinene secondary
- 1020 organic aerosol, Proc. Natl. Acad. Sci. U. S. A., 112, 14168-14173,
- 1021 10.1073/pnas.1517742112, 2015b.





1038	Table 1: Experimenta	l conditions and	l results for the $\alpha$	-pinene ozonol	ysis experiments	
------	----------------------	------------------	----------------------------	----------------	------------------	--

	I			2	1
	Experiment	Initial Seed	Initial	$\Delta M_0^{\ b}$	SOA Mass
		Surface Area	[α-pinene] <sup>a</sup>	$(\mu g m^{-3})$	Yield <sup>c</sup> (%)
		$(\mu m^2 cm^{-3})$	$(\mu g m^{-3})$		
	100 ppb O <sub>3</sub> nucleation	0	290.2±23.2	$62.0\pm1.2^{d}$	22.6±1.9
	100 ppb O <sub>3</sub> low AS	1130	280.5±22.4	$63.0 \pm 0.8^{d}$	23.3±1.9
	100 ppb O <sub>3</sub> high AS	2700	238.7±19.1	$50.6 \pm 1.6^{d}$	23.3±1.9
	500 ppb O <sub>3</sub> nucleation	0	274.4±21.9	$87.3 \pm 0.3^{e}$	31.8±2.5
	500 ppb O <sub>3</sub> low AS	1300	264.9±21.2	$75.7 \pm 0.6^{e}$	28.6±2.3
	500 ppb O <sub>3</sub> high AS	2720	236.1±18.9	$66.3 \pm 1.9^{e}$	28.1±2.4
20	30	· · · · 1 · · · · · · 1	1 1 411/1	1 •	4 1 41 500

<sup>a</sup>Concentration of  $\alpha$ -pinene injected into the chamber. All the  $\alpha$ -pinene reacted in the 500 ppb O<sub>3</sub> experiments, but not the 100 ppb O<sub>3</sub> experiments.

1041 <sup>b</sup>Uncertainties in the peak SOA mass concentration ( $\Delta M_0$ ) are calculated from one 1042 standard deviation of the aerosol volume as measured by the scanning mobility particle 1043 sizer.

1044 <sup>c</sup>SOA mass yields at peak SOA growth are reported.

<sup>d</sup>The SOA mass concentration is calculated using the density =  $1.39 \text{ g cm}^{-3}$  obtained from the 100 ppb O<sub>3</sub> nucleation experiment.

<sup>1040</sup> <sup>e</sup>The SOA mass concentration is calculated using the density =  $1.37 \text{ g cm}^{-3}$  obtained from 1048 the 500 ppb O<sub>3</sub> nucleation experiment.

1049

1050

1051

1052

1053

1054

1055

1056

1057

1058

1059





# **Table 2:** Coupled vapor-particle dynamics model parameters

Parameter	Definition	Value
$\alpha_p$	Vapor-particle mass accommodation coefficient	1
$\alpha_w$	Vapor-wall mass accommodation coefficient	10-6
$ au_{olig}$	Timescale of oligomerization	4 h
$C^*$	Saturation vapor pressures and branching ratios of	$[0.57 (>10^3), 0.35$
	oxidation products	$(10^2), 0.04 (10), 0.015$
		(1) and 0.025 (0.1)]
$D_i$	Gas-phase molecular diffusivity	$3 \times 10^{-6} \text{ m}^2 \text{ s}^{-1}$
A/V	Surface area-to-volume ratio of the chamber	$2.5 \text{ m}^{-1}$
$C_w$	Equivalent organic mass concentration in the wall	10 mg m <sup>-3</sup>
ke	Eddy diffusion coefficient	$0.03 \text{ s}^{-1}$
$M_i$	Molecular weight of the diffusing gas-phase molecule i	168, 184, 192, 200 and
		216 g mole <sup>-1</sup>
Minit	Initially absorbing organic material in seed aerosol	0.01 μg m <sup>-3</sup>
Р	Pressure	$1 \times 10^5$ Pa
Т	Temperature	298 K
$\rho_{seed}$	Density of inorganic seed	1700 kg m <sup>-3</sup>
$\rho_{org}$	Density of organic material on seed particle	1300 kg m <sup>-3</sup>







1069

1070 **Figure 1:** Particle wall deposition coefficients ( $\beta_i$ ) measured during the low AS-seed 1071 only and high AS-seed only experiments in GTEC. Also shown are the particle wall 1072 deposition coefficients (labeled "Other") measured in previous routine monthly AS-seed 1073 only experiments in the chamber. These previous routine monthly AS-seed only 1074 experiments were performed using either a 0.008 M AS or a 0.1 M AS solution.







1076

1077 Figure 2: Reaction profiles of the  $\alpha$ -pinene ozonolysis experiments. Panels (a), (b) and 1078 (c) show results from the nucleation, low AS and high AS 100 ppb O<sub>3</sub> experiments, 1079 respectively. Panels (d), (e) and (f) show results from the nucleation, low AS and high AS 500 ppb O<sub>3</sub> experiments, respectively. As explained in the main text, the SOA mass 1080 concentrations ( $\Delta M_o$ ) for the nucleation and low AS experiments are obtained using the 1081 1082 particle wall rates obtained from the low AS-seed only experiments, while the SOA mass 1083 concentrations ( $\Delta M_0$ ) for the high AS-seed experiments are obtained using the particle 1084 wall rates obtained from the high AS-seed only experiments.







1086

1087Figure 3: Time-dependent SOA growth curves for α-pinene ozonolysis. Panels (a) and1088(b) show 10 min-averaged results from the 100 ppb and 500 ppb  $O_3$  experiments,1089respectively. Only SOA growth data up to the point of SOA peak growth are shown.







1091

1092 **Figure 4:** 10 min-averaged SOA mass yields over the course of an  $\alpha$ -pinene ozonolysis 1093 experiment as a function of initial total AS seed surface area concentration for the (a) 100 1094 ppb  $O_3$  experiments, and (b) 500 ppb  $O_3$  experiments. Symbol color indicates the SOA mass concentration and symbol size indicates the time after O<sub>3</sub> is injected into the 1095 1096 chamber. The  $\times$  symbols are the SOA mass yields at peak SOA growth obtained from 1097 the experimental data. The y-axis error bars represent the uncertainty in the SOA mass 1098 yield at peak SOA growth, which originates from the  $\alpha$ -pinene injection and the aerosol 1099 volume concentration measured by the SMPS at peak SOA growth (one standard 1100 deviation).







1104 Figure 5: Comparison of SOA mass yields obtained in this study to those of previous 1105 dark  $\alpha$ -pinene ozonolysis studies (Table S2). The SOA mass yields and concentrations of majority of these previous studies (Hoffmann et al., 1997; Griffin et al., 1999; Cocker et 1106 1107 al., 2001b; Gao et al., 2004; Presto et al., 2005; Presto et al. 2006; Pathak et al., 2007b; 1108 Song et al., 2007) were previously compiled by Shilling et al (2008). Similar to Shilling 1109 et al. (2008), all the data shown here (including those reported in this study) have been adjusted using an organic density of 1.0 g cm<sup>-3</sup>, and to 298 K using a temperature 1110 1111 correction of 1.6 % per K, as recommended by Pathak et al. (2007b) to facilitate easier 1112 comparison among the different studies.









1114 **Figure 6:** SOA mass yields at peak SOA growth as a function of both the seed surface 1115 area and  $O_3$  concentration for  $\alpha_p = 1$ , 0.01, and 0.001. The SOA mass yields at peak SOA 1116 growth are indicated by colors and contours. Note that the color bars for panels (a), (b) 1117 and (c) have different SOA mass yield ranges. Simulations were carried out using the 1118 branching ratios, oligomerization rate, and vapor wall deposition rate parameters obtained 1119 in this study. The initial  $\alpha$ -pinene concentration was set to 50 ppb, and a fixed  $O_3$ 1120 concentration was used in place of a linear injection.