1	Influence of Seed Aerosol Surface Area and Oxidation Rate on Vapor-Wall
2	Deposition and SOA Mass Yields: A case study with α-pinene Ozonolysis
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Abstract

14

15 Laboratory chambers, invaluable in atmospheric chemistry and aerosol formation studies, 16 are subject to particle and vapor wall deposition, processes that need to be accounted for 17 in order to accurately determine secondary organic aerosol (SOA) mass yields. Although 18 particle wall deposition is reasonably well understood and usually accounted for, vapor 19 wall deposition is less so. The effects of vapor wall deposition on SOA mass yields in 20 chamber experiments can be constrained experimentally by increasing the seed aerosol 21 surface area to promote the preferential condensation of SOA-forming vapors onto seed aerosol. Here, we study the influence of seed aerosol surface area and oxidation rate on 22 23 SOA formation in  $\alpha$ -pinene ozonolysis. The observations are analyzed using a coupled 24 vapor-particle dynamics model to interpret the roles of gas-particle partitioning (quasi-25 equilibrium vs. kinetically-limited SOA growth) and  $\alpha$ -pinene oxidation rate in 26 influencing vapor wall deposition. We find that the SOA growth rate and mass yields are 27 independent of seed surface area within the range of seed surface area concentrations used in this study. This behavior arises when the condensation of SOA-forming vapors is 28 29 dominated by quasi-equilibrium growth. Faster  $\alpha$ -pinene oxidation rates and higher SOA 30 mass yields are observed at increasing  $O_3$  concentrations for the same initial  $\alpha$ -pinene 31 concentration. When the  $\alpha$ -pinene oxidation rate increases relative to vapor wall 32 deposition, rapidly produced SOA-forming oxidation products condense more readily 33 onto seed aerosol particles, resulting in higher SOA mass yields. Our results indicate that 34 the extent to which vapor wall deposition affects SOA mass yields depends on the 35 particular VOC system, and can be mitigated through the use of excess oxidant36 concentrations.

### **1. Introduction**

38 Secondary organic aerosol (SOA), formed from the oxidation of volatile and 39 intermediate volatility organic compounds (VOCs and IVOCs), contributes a significant 40 fraction of the global organic aerosol burden (Kanakidou et al., 2005; Hallquist et al., 41 2009; Tsigaridis et al., 2014). SOA formation studies, which are typically conducted in 42 laboratory chambers in the presence of seed aerosol particles, provide fundamental data 43 that can be used to predict the rate of atmospheric SOA formation. An essential 44 parameter of interest in laboratory chamber studies is the SOA mass yield (Y), which is 45 defined as the ratio of mass concentration of SOA formed to mass concentration of parent 46 hydrocarbon reacted (( $\Delta$ HC), Y =  $\Delta$ M<sub>o</sub>/ $\Delta$ HC (Odum et al., 1996; Odum et al., 1997a; 47 Odum et al., 1997b)). The measured SOA mass yields can subsequently be applied in 48 atmospheric models to predict regional and global organic aerosol burdens. In order to 49 obtain accurate SOA mass yields from the evolving aerosol size distribution in chamber 50 experiments, the loss of both particles and vapors to the chamber walls needs to be 51 accurately accounted for (Crump and Seinfeld, 1981; McMurry and Grosjean, 1985; 52 McMurry and Rader, 1985; Cocker et al., 2001a; Weitkamp et al., 2007; Pierce et al., 53 2008; Hildebrandt et al., 2009; Loza et al., 2010; Matsunaga and Ziemann, 2010; Loza et 54 al., 2012; Kokkola et al., 2014; McVay et al., 2014; Yeh and Ziemann, 2014; Zhang et 55 al., 2014; Yeh and Ziemann, 2015; Zhang et al., 2015a; La et al., 2016; Ye et al., 2016).

56 The mechanisms by which particles in chambers deposit on chamber walls are 57 reasonably well understood. Particles are transported to the boundary layer on the 58 chamber walls via diffusion, gravitational settling, and electrostatic forces (Crump and 59 Seinfeld, 1981; McMurry and Grosjean, 1985; McMurry and Rader, 1985; Pierce et al., 60 2008). The rate at which particles are transported to the edge of the boundary layer is 61 dictated primarily by mixing conditions in the chamber. An effective approach for 62 characterizing particle wall loss involves measuring the size-dependent wall loss rates of 63 polydisperse inert seed aerosol (e.g. ammonium sulfate particles) injected into the 64 chamber during seed-only experiments (Keywood et al., 2004; Pierce et al., 2008). The

observed particle number concentration decay in each size bin is then fitted to a firstorder exponential decay from which the first-order wall loss coefficients are determined as a function of particle size. These wall loss coefficients are subsequently used to correct for size-dependent particle wall loss in actual SOA formation experiments. Determination of particle wall loss coefficients may be complicated if coagulation is significant. Particle dynamics models can be used to correct particle wall loss coefficients for coagulation.

71 Vapor-wall deposition mechanisms in chambers are not as well understood or 72 accounted for as those for particles. The degree to which SOA-forming vapors deposit 73 onto chamber walls is governed by the rate at which these gas-phase organic molecules 74 are transported to the walls, the strength of adherence of the organic molecule to the wall, 75 and the extent of reversible vapor-wall partitioning (Loza et al., 2010; Matsunaga and 76 Ziemann, 2010; Zhang et al., 2015a). For example, Loza et al. (2010) showed that the 77 loss of 2,3-epoxy-1,4-butanediol, an isoprene oxidation product analogue, to walls in the 78 Caltech chamber was essentially irreversible on short time scales but became reversible 79 on longer time scales. In contrast, glyoxal, a common isoprene oxidation product, 80 exhibited reversible vapor-wall partitioning over all time scales. Recent studies show that 81 SOA mass yields measured in chamber experiments can be significantly underestimated 82 due to wall deposition of SOA-forming vapors that would otherwise contribute to SOA growth (McVay et al., 2014; Zhang et al., 2014; La et al., 2016). Zhang et al. (2014) 83 84 found that chamber-derived SOA mass yields from toluene photooxidation may be 85 underestimated by as much as a factor of four as a result of vapor wall loss. 86 Consequently, the use of underestimated chamber-derived SOA mass yields in 87 atmospheric models will lead to the underprediction of ambient SOA mass concentrations 88 (Cappa et al., 2016).

For the toluene photooxidation system, Zhang et al. (2014) showed that the measured SOA mass yields increased with increasing seed aerosol surface area, demonstrating that increasing the seed-to-chamber surface area ratio promoted the condensation of SOA-forming vapors onto seed aerosol particles. However, increasing the seed aerosol surface area to promote condensation of SOA-forming vapors onto seed aerosol particles may not be effective in all VOC oxidation systems. A modeling study by

95 McVay et al. (2014) showed that the SOA mass yield depends on seed aerosol surface 96 area only in cases where the condensation of SOA-forming vapors onto seed aerosol 97 particles is kinetically limited (i.e., the timescale for gas-particle equilibrium is 98 competitive with or greater than the timescale for reaction and vapor-wall deposition). In 99 addition to the seed aerosol surface area, VOC oxidation rate may also play an important 100 role in the effect of vapor wall loss on SOA formation. Ng et al. (2007) showed that the 101 SOA mass yields from *m*-xylene photooxidation are dependent on the oxidation rate, with 102 higher OH concentrations (hence faster oxidation rates) resulting in higher SOA mass 103 vields. It was suggested that the "oxidation rate effect" could arise as a result of 104 competition between growing particles and chamber walls for condensable VOC 105 oxidation products (Ng et al., 2007). However, McVay et al. (2016) reported similar SOA 106 growth at low and high OH concentrations in  $\alpha$ -pinene photooxidation. Taken together, 107 these studies show the importance of understanding how gas-particle partitioning and 108 VOC oxidation rate impact vapor-wall deposition and SOA mass yields in laboratory 109 chamber experiments.

110 In this study, we examine the influence of seed aerosol surface area and oxidation 111 rate on SOA formation in  $\alpha$ -pinene ozonolysis chamber experiments.  $\alpha$ -pinene is the most abundant monoterpene, with global emissions estimated to be  $\sim 66 \text{ Tg vr}^{-1}$  (Guenther et 112 al., 2012). Ozonolysis is the major atmospheric oxidation pathway of  $\alpha$ -pinene, and is 113 estimated to account for reaction of ~46 % of emitted  $\alpha$ -pinene (Griffin et al., 1999; 114 115 Capouet et al., 2008).  $\alpha$ -pinene ozonolysis, a major source of atmospheric SOA on both 116 regional and global scales (Kanakidou et al., 2005; Hallquist et al., 2009; Carlton et al., 117 2010; Pye et al., 2010), has been the subject of numerous studies (Hoffmann et al., 1997; Griffin et al., 1999; Cocker et al., 2001b; Gao et al., 2004; Presto et al., 2005; Presto and 118 119 Donahue, 2006; Pathak et al., 2007a; Pathak et al., 2007b; Song et al., 2007; Shilling et 120 al., 2008; Henry et al., 2012; Ehn et al., 2014; Kristensen et al., 2014; Zhang et al., 121 2015b). Here, we measure the  $\alpha$ -pinene SOA mass yield as a function of seed aerosol surface area concentration (0 to 3000  $\mu$ m<sup>2</sup> cm<sup>-3</sup>) and O<sub>3</sub> mixing ratio (100 vs. 500 ppb). 122 123 These results are analyzed using a coupled vapor-particle dynamics model to evaluate the 124 roles of gas-particle partitioning and VOC oxidation rate in influencing vapor-wall 125 deposition effects on the measured SOA mass yields.

# 126 **2. Experimental**

# 127 **2.1. Dark** α-pinene ozonolysis experiments

128 Experiments were conducted in the Georgia Tech Environmental Chamber (GTEC) facility. Details of the dual chamber facility are provided elsewhere (Boyd et al., 129 130 2015). Only one FEP Teflon chamber (volume 13  $m^3$ ) was used for the entirety of this 131 study. Before each experiment, the chamber was flushed with dried, purified air for at least 36 h until the aerosol number concentration was  $< 30 \text{ cm}^{-3}$ . All experiments were 132 conducted under dry conditions (< 5 % RH) at room temperature (25 °C). NO<sub>x</sub> mixing 133 134 ratios in these experiments were < 1 ppb. Experimental conditions are summarized in 135 Table 1.

136 22 ppm of cyclohexane (Sigma Aldrich,  $\geq$  99.9 %) was first injected into the chamber to act as an OH scavenger (~440 times the initial  $\alpha$ -pinene concentration). Based 137 138 on the concentrations of cyclohexane and  $\alpha$ -pinene injected into the chamber, the reaction 139 rate of OH with cyclohexane is ~60 times greater than that with  $\alpha$ -pinene. After the 140 cyclohexane concentration had stabilized in the chamber for 30 min, a known concentration (~50 ppb in all experiments) of  $\alpha$ -pinene (Sigma Aldrich, > 99 %) was 141 142 injected into the chamber, followed by inorganic seed aerosol via atomization of an 143 aqueous ammonium sulfate (AS) solution (in seeded experiments). To vary the seed 144 aerosol surface area, different concentrations of AS solutions were used to generate seed 145 aerosol particles in the seeded experiments. In the "low AS" experiments, a 0.015 M AS 146 solution was used to generate seed particles, and the resulting initial total AS seed surface area concentration was  $\sim 1000 \text{ }\mu\text{m}^2 \text{ cm}^{-3}$ . In the "high AS" experiments, a 0.05 M AS 147 solution was used to generate seed aerosol particles, and the resulting initial total AS seed 148 surface area concentration was  $\sim 3000 \text{ } \text{ } \text{ } \text{m}^{-3}$ . In selected experiments, no seed aerosol 149 150 particles were introduced into the chamber and SOA was formed via nucleation. After the 151 seed aerosol concentration in the chamber stabilized,  $O_3$  (100 or 500 ppb), which was 152 generated by passing purified air into a photochemical cell (Jelight 610), was introduced 153 into the chamber. The start of  $O_3$  injection into the chamber marked the beginning of the reaction (i.e., reaction time = 0 min). The injected  $\alpha$ -pinene:O<sub>3</sub> molar ratio was 154 155 approximately 1:2 and 1:10 in the 100 and 500 ppb O<sub>3</sub> experiments, respectively. O<sub>3</sub> was

156 injected into the chamber for 13.5 and 54.25 min in the 100 and 500 ppb O<sub>3</sub> experiments, 157 respectively, to achieve the desired  $O_3$  concentrations. Approximately 11 % and 98 % of 158 the initial  $\alpha$ -pinene had reacted when O<sub>3</sub> injection was completed in the 100 and 500 ppb 159  $O_3$  experiments, respectively. In the GTEC chamber,  $\alpha$ -pinene closest to the  $O_3$  injection 160 port likely reacted first in the  $\alpha$ -pinene ozonolysis experiments. The O<sub>3</sub> injection times 161 were established in separate experiments in which only  $O_3$  was injected into the chamber. 162 Based on the  $O_3$  time series traces in the  $O_3$ -only experiments, the  $O_3$  mixing timescale 163 was estimated to be  $\sim 12$  min for all experiments.

164 The  $\alpha$ -pinene and O<sub>3</sub> concentrations were measured by a Gas Chromatograph-165 Flame Ionization Detector (GC-FID, Agilent 7890A) and O<sub>3</sub> monitor (Teledyne T400), 166 respectively. GC-FID measurements were taken 12 min apart. A High Resolution Time-167 of-Flight Aerosol Mass Spectrometer (HR-ToF-AMS, Aerodyne Research Inc.) was used 168 to measure the aerosol elemental composition (DeCarlo et al., 2006; Canagaratna et al., 169 2015). Details on the operation of the HR-ToF-AMS and its data analysis are described 170 elsewhere (Canagaratna et al., 2015). Aerosol size distributions, number and volume 171 concentrations were measured by a Scanning Mobility Particle Sizer (SMPS, TSI), which 172 consists of a Differential Mobility Analyzer (DMA, TSI 3081) and a Condensation 173 Particle Counter (CPC, TSI 3775). For nucleation and low AS experiments, the measured 174 aerosol size range was set to 14 to 686 nm diameter. For high AS experiments, the 175 measured aerosol size range was set to 17 to 983 nm. Prior checks were made to confirm 176 that no particles larger than 686 nm were detected in the nucleation and low AS 177 experiments. The SOA mass concentrations reported in this study were measured using 178 the SMPS. The SOA density was calculated from the ratio of the aerosol size 179 distributions measured by the HR-ToF-AMS and the SMPS during nucleation 180 experiments (DeCarlo et al., 2004; Bahreini et al., 2005).

## 181 **2.2. Particle wall deposition correction**

Particle wall deposition needs to be accounted for to determine the SOA mass concentration in the chamber. Two limiting assumptions have traditionally been made regarding interactions between particles deposited on the chamber walls and suspended vapors when accounting for particle wall loss in the computation of SOA mass yields

186 (Weitkamp et al., 2007; Hildebrandt et al., 2009; Loza et al., 2012; Zhang et al., 2014). 187 The first case assumes that particles deposited on the walls cease to interact with 188 suspended vapors, and therefore the SOA mass present on these deposited particles does 189 not change after deposition (Loza et al., 2012; Zhang et al., 2014). Adding the SOA mass 190 present on these deposited particles to that present on the suspended particles provides a 191 lower bound of the total SOA mass concentration. In the second case, it is assumed that 192 particles deposited on the walls continue to interact with suspended vapors as if these 193 particles had remained suspended, and therefore the SOA mass present on these 194 deposited particles increases at the same rate as those suspended (Hildebrandt et al., 195 2009; Weitkamp et al., 2007). Thus, this case provides an upper bound of the total SOA 196 mass concentration due to the additional uptake of suspended vapors to wall-deposited 197 particles. However, it must be kept in mind that the calculated SOA mass concentration 198 can be underestimated even in the upper bound case since the calculation accounts 199 neither for differences in the vapor-particle and vapor-wall interaction and transport 200 timescales nor for the significantly larger amount of absorbing mass of the chamber walls 201 (relative to the deposited particles) for suspended vapors (McVay et al., 2014; Zhang et 202 al., 2014; McVay et al., 2016).

In this study, we calculate SOA mass yields using the lower bound of the total SOA mass concentration obtained from SMPS measurements, which has been described in detail previously (Loza et al., 2012), and will be reviewed briefly here. For each particle size bin *i* at each time increment  $\Delta t$ , the particle number distribution deposited on the wall  $(n_{w,i,i})$  is:

208 
$$n_{w,i,j} = n_{s,i,j} \times (1 - \exp(-\beta_i \Delta t))$$

where  $n_{s,i,j}$  is the suspended particle number distribution in particle size bin *i* at time step *j*,  $\Delta t$  is the difference between time step *j* and time step *j* + 1, and  $\beta_i$  is the size-dependent first-order exponential wall loss rate obtained from seed-only experiments. As we describe subsequently,  $\beta_i$  may be measured directly during seed-only experiments or may be corrected for the influence of coagulation using a particle dynamics model. The

(1)

particle wall loss corrected number distribution  $(n_{total,i,j})$  is obtained from the sum of the particle number distribution of deposited particles  $(n_{w,i,j})$  and suspended particles  $(n_{s,i,j})$ :

216 
$$n_{total,i,j} = n_{s,i,j} + n_{w,i,j}$$
(2)

Assuming spherical particles, the particle wall loss corrected volume concentration  $(V_{total,j})$  is:

219 
$$V_{total,j} = \sum_{i=1}^{m} \frac{n_{total,i,j}}{D_{p,i} \ln 10} \times (D_{p,i+} - D_{p,i-}) \times \frac{\pi}{6} D_{p,i}^{3}$$
(3)

where *m* is the number of particle size bins,  $D_{p,i^+}$  and  $D_{p,i^-}$  are the upper and lower limits for size bin *i*, respectively, and  $D_{p,i}$  is the median particle diameter for size bin *i*. The term  $D_{p,i}$ ln10 is needed to convert from a lognormal distribution. Figures S1-S4 and Table S1 show results from the particle wall loss correction. To calculate the SOA mass concentration ( $\Delta M_{o,j}$ ), the SOA density ( $\rho_{org}$ ) is multiplied by the difference of the particle wall loss corrected volume concentration ( $V_{total,j}$ ) and the initial seed volume concentration ( $V_{seed}$ ):

227 
$$\Delta M_{o,j} = \rho_{org} \times (V_{total,j} - V_{seed})$$
(4)

The measured densities of the  $\alpha$ -pinene SOA are 1.39 and 1.37 g cm<sup>-3</sup> for the 100 and 500 ppb O<sub>3</sub> experiments, respectively, and are within the range (i.e., 1.19 to 1.52 g cm<sup>-3</sup>) reported in previous  $\alpha$ -pinene ozonolysis studies (Bahreini et al., 2005; Kostenidou et al., 2007; Song et al., 2007; Shilling et al., 2009).

# 232 **3. Vapor-particle dynamics model**

A coupled vapor-particle dynamics model is used to evaluate the influence of seed aerosol surface area and oxidation rate on SOA formation in the  $\alpha$ -pinene ozonolysis chamber experiments. This model is similar to that used in McVay et al. (2014), and will be briefly described here. Parameters from the experimental data (temperature, pressure, initial  $\alpha$ -pinene concentration) are used as model inputs. The initial size distribution is set to that measured by the SMPS, with the exception of the two nucleation experiments. Because nucleation is not explicitly simulated, an approximation is used in which the 240 smallest diameter bin is initialized with the total number of particles measured at the end 241 of the experiment (see Table S1). In each simulation, the decay of  $\alpha$ -pinene, the 242 consumption of O<sub>3</sub>, the SOA mass concentration, and the SOA mass yield are calculated throughout the duration of the experiment. We assume a linear injection rate of O<sub>3</sub> based 243 244 on the time required to inject the desired  $O_3$  concentration. For example,  $O_3$  is injected at a rate of 500/54.25 ppb min<sup>-1</sup> for the first 54.25 min during the 500 ppb  $O_3$  experiments. 245 O<sub>3</sub> simultaneously decays by reaction with  $\alpha$ -pinene at a rate constant of 9.4 × 10<sup>-17</sup> cm<sup>3</sup> 246 molec.<sup>-1</sup> s<sup>-1</sup> (Saunders et al., 2003). Modeled  $O_3$  and  $\alpha$ -pinene concentrations are 247 compared with observed concentrations in Fig. S5. The good fit of modeled and observed 248 249  $O_3$  and  $\alpha$ -pinene concentrations indicates that our representation of  $O_3$  is appropriate. The 250  $O_3+\alpha$ -pinene reaction is assumed to occur in a well-mixed chamber and produces 5 251 classes of first-generation products, which are grouped according to mass saturation concentrations, similar to the volatility basis set (Donahue et al., 2006):  $>10^3$  (assumed to 252 be completely volatile), 10<sup>2</sup>, 10, 1 and 0.1 µg m<sup>-3</sup>. Branching ratios between these 253 products are optimized to fit the experimental data. These branching ratios cannot be 254 255 compared directly to previously reported VBS parameters for  $\alpha$ -pinene ozonolysis (e.g. 256 Henry et al. (2012)) since VBS parameters are typically mass-based, while the branching 257 ratios in the model are mole-based. Furthermore, the branching ratios here account for the 258 influence of vapor wall deposition, while typical VBS parameters do not. We assume that these 5 classes of products have molecular weights 168, 184, 192, 200 and 216 g mole<sup>-1</sup> 259 260 based on the group contribution method (Donahue et al., 2011). The first-generation 261 products are assumed not to undergo further reaction with O<sub>3</sub> upon formation.

262 The aerosol dynamics in the chamber obey the aerosol general dynamic equation263 (Seinfeld and Pandis, 2006):

264 
$$\left(\frac{\partial n(D_p,t)}{\partial t}\right) = \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{coag}} + \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{cond/evap}} + \left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{wall loss}}$$
(5)

Coagulation is not considered in the present model; we address the potential impact of coagulation later in the paper. The change in particle number distribution due to particle wall loss is:

268 
$$\left(\frac{\partial n(D_p,t)}{\partial t}\right)_{\text{wall loss}} = -\beta_j(D_p)n(D_p,t)$$
(6)

where, as noted in section 2.2,  $\beta_j(D_p)$  is the size-dependent first-order wall loss rate coefficient obtained from fitting seed-only experiments. The rate at which vapor condenses onto a spherical aerosol particle is:

272 
$$J_{i} = 2\pi D_{p} D_{i} (G_{i} - G_{i}^{eq}) F_{FS}$$
(7)

where  $G_i$  is the concentration of gas-phase species *i*,  $G_i^{eq}$  is the saturation concentration of gas-phase species *i*,  $D_i$  is the gas-phase molecular diffusivity (assumed to be 3 × 10<sup>-6</sup> m<sup>2</sup> s<sup>-1</sup> (McVay et al., 2014)), and  $F_{FS}$  is the Fuchs-Sutugin correction for non-continuum gas-phase diffusion:

277 
$$F_{FS} = \frac{0.75\alpha_p (1 + \text{Kn})}{\text{Kn}^2 + \text{Kn} + 0.283\text{Kn}\alpha_p + 0.75\alpha_p}$$
(8)

where  $\alpha_p$  is the vapor-particle mass accommodation coefficient, and Kn is the Knudsen number, Kn =  $2\lambda_{AB}/D_p$ . The vapor-particle mass accommodation coefficient accounts for any resistance to vapor molecule uptake at the particle surface (e.g. surface accommodation and particle-phase diffusion limitations).  $\lambda_{AB}$  is the mean free path of the gas-phase species, which is:

283 
$$\lambda_{AB} = 3D_i \times \sqrt{\frac{\pi M_i}{8RT}}$$
(9)

where *R* is the ideal gas constant, *T* is the temperature, and  $M_i$  is the molecular weight of diffusing gas-phase molecule *i*. For each particle size bin, Eqs. 7-9 are used to compute the flux of each gas-phase species to and from an aerosol particle, scaled by the particle number concentration in the size bin. The net rate of change for each gas-phase species due to evaporation or condensation is obtained from the total flux summed over all the particle size bins. 290  $G_i^{eq}$  varies for each particle size bin because it depends on the mass concentration 291 of species *i* and the total organic mass concentration in the size bin:

292 
$$G_i^{eq} = \frac{A_i C_i^*}{\sum_k A_k + M_{init}}$$
(10)

where  $A_i$  is the concentration of species *i* in the particle phase,  $C_i^*$  is the saturation concentration of species *i*,  $\sum_k A_k$  is the sum of all the species concentration in the particle phase, and  $M_{init}$  is the mass concentration of any absorbing organic material initially present in the seed aerosol. To avoid numerical errors in Eq. 10 at the first time step,  $M_{init}$ is set to 0.01 µg m<sup>-3</sup>.

298 The oxidation products of  $\alpha$ -pinene ozonolysis are assumed to be subject to 299 vapor-wall deposition, which is simulated using a first-order wall-loss coefficient 300 (McMurry and Grosjean, 1985):

301
$$k_{wall,on} = \left(\frac{A}{V}\right) \frac{\frac{\alpha_{wall}c}{4}}{1 + \frac{\pi}{2}\left(\frac{\alpha_{wall}c}{4\sqrt{k_e D_i}}\right)}$$
(11)

where A/V is the surface area-to-volume ratio of the chamber (estimated to be 2.5 m<sup>-1</sup>),  $\alpha_{wall}$  is the vapor-wall mass accommodation coefficient, and  $k_e$  is the eddy diffusion coefficient that describes mixing conditions in the chamber. Based on the measured sizedependent particle wall loss rates (method is described in Zhang et al. (2014)),  $k_e$  is estimated to be 0.03 s<sup>-1</sup> for the GTEC chamber. Vapor-wall deposition is assumed to be reversible, and the rate constant of vapor desorption from the chamber walls is:

$$k_{wall,off} = \frac{k_{wall,on}}{K_W C_W} = k_{wall,on} \left(\frac{C_i^* M_w \gamma_w}{C_W M_p \gamma_p}\right)$$
(12)

309 where  $C_w$  is the equivalent organic mass concentration in the wall (designated to treat 310 gas-wall partitioning in terms of gas-particle partitioning theory and not necessarily

11

311 representative of a physical layer of organic concentration on the wall (Matsunaga and 312 Ziemann, 2010)),  $K_w$  is the gas-wall partitioning coefficient,  $M_w$  is the effective molecular 313 weight of the wall material,  $\gamma_w$  is the activity coefficient of the species in the wall layer, 314  $M_p$  is the average molecular weight of organic species in the particle, and  $\gamma_p$  is the activity coefficient of the species in the particle. For simplicity, we assume that  $M_w = M_p$  and  $\gamma_w =$ 315  $\gamma_p$ .  $C_w$  is set to 10 mg m<sup>-3</sup> based on previous inferences by Matsunaga and Ziemann 316 (2010). Sensitivity studies (not shown) show no change in model predictions when 317 318 varying  $C_w$  above  $C_w = 0.1 \text{ mg m}^{-3}$ .

319 In the initial version of the model, after all the  $\alpha$ -pinene is consumed, vapor wall 320 deposition was assumed to continue to deplete the gas-phase oxidation products and 321 aerosol mass evaporates to maintain gas-particle equilibrium. SOA evaporation was not 322 observed experimentally (i.e., the SOA mass concentration does not decrease 323 significantly over time after peak SOA growth has been achieved in these chamber 324 experiments (Fig. 2)). In order to represent these observations in the model, a first-order, 325 particle-phase reaction is introduced by which aerosol species are converted into non-326 volatile absorbing organic mass with a timescale of  $\tau_{olig}$ . This mechanism (which is not 327 included in the model used in McVay et al. (2014)) is similar to that used by the 328 sequential equilibrium partitioning model, in which aerosol is converted from an 329 absorbing to non-absorbing, non-volatile phase in order to explain the inhibited diffusion 330 and evaporation observed in  $\alpha$ -pinene ozonolysis SOA (Cappa and Wilson, 2011). 331 Although we assume here that the converted non-volatile aerosol mass still participates in 332 partitioning, either mechanism invokes a particle-phase process to retard SOA 333 evaporation.

In order to determine the parameters for  $\alpha_w$ ,  $\alpha_p$ ,  $\tau_{olig}$  and the branching ratios between the oxidation products that provide the best fit to measured SOA data, the parameter space was discretized and all possible combinations of parameters were simulated, following Karnezi et al. (2014). In order to restrict the number of combinations required, only parameter values judged to be physically realistic were chosen. Because the branching ratios in this model are mole-based, they must sum to one; therefore only combinations of parameters summing to one were allowed. The discretization is shown in Table S2 and results in roughly 10,000 different combinations
of parameters. All six experiments were simulated with each parameter combination, and
simulations were run using GNU Parallel (Tange 2011). For each combination of
parameters, the percentage error was calculated from equation 10 of Karnezi et al.
(2014):

346 
$$E_i = \frac{100}{n} \sqrt{\sum_{i}^{n} (Moa_{i,guess} - Moa_{i,meas})^2}$$
(13)

where  $Moa_{i,guess}$  is the model-predicted SOA mass concentration at a particular timestep *i* for one of the experiments,  $Moa_{i,meas}$  is the measured SOA mass concentration at a particular timestep *i* for one of the experiments, and *n* is the number of timesteps summed over all experiments. The best-fit combination of parameters is defined as the combination of parameters with the lowest percentage error. This lowest-error combination of parameters was compared to the "best estimate" parameters determined from the inverse error weighting factor of Karnezi et al. (2014):

354 
$$\bar{x} = \frac{\sum_{j}^{N} \left[ x_{j} \frac{1}{E_{j}} \right]}{\sum_{j}^{N} \frac{1}{E_{j}}}$$
(14)

where  $x_j$  is a value of one of the parameters ( $\alpha_w$ ,  $\alpha_p$ ,  $\tau_{olig}$  or a branching ratio between the oxidation products), with N different possible combinations of parameters, and  $E_j$  is the percent error for that particular combination of parameters. The lowest-error combination of parameters and the Karnezi et al. (2014) best estimate parameters are both reported, but the lowest-error combination of parameters resulted in a lower percentage error than the Karnezi et al. (2014) best estimate parameters. The lowest-error combination of parameters is used for the modeling analysis.

### **362 4. Results**

Red and blue solid lines in Fig. 1 shows the size-dependent particle wall deposition coefficients measured in the low AS-seed only and high AS- seed-only deposition experiments. In these measurements, we assume that the number concentration is low enough such that the effect of coagulation is small and only particle wall 367 deposition affects the particle size distribution, thus allowing for the direct measurement 368 of size-dependent particle wall deposition coefficients. The initial total AS seed surface 369 area concentration in the low AS-seed only and high AS-seed only experiments (which 370 are conducted using 0.015 M AS and 0.05 M AS solutions, respectively) are similar to those used in the  $\alpha$ -pinene ozonolysis experiments (i.e., ~1000 and ~3000  $\mu$ m<sup>2</sup> cm<sup>-3</sup>, 371 respectively). As shown in Fig. 1, the measured particle wall deposition coefficients from 372 373 the low AS-seed only and high AS-seed only experiments generally fall within the range 374 of those measured in routine monthly AS-seed only experiments conducted in the 375 chamber. Figure 1 also shows the size-dependent particle wall deposition coefficients 376 corrected for coagulation, shown using dashed lines, which are obtained as described in 377 Pierce et al., 2008 using the data from the low AS-seed only and high AS-seed only 378 experiments. A comprehensive description of the relationship between coagulation and 379 particle wall deposition will be provided in a forthcoming publication. Briefly, as 380 described in Pierce et al. (2008), the coagulation-corrected particle wall loss coefficients 381 are determined by simulating the decay of the initial size distribution due to coagulation 382 and then attributing the difference in this decay and the observed decay to particle wall 383 loss. For both the measured and coagulation-corrected particle wall deposition 384 coefficients, the minimum coefficient for the low AS-seed only experiment is different 385 from that of the high AS-seed only experiments. The cause of this difference is currently 386 under investigation but may be due in part to uncertainties arising from the low particle 387 number concentrations for the larger particles in the low AS-seed only experiment. To 388 study how coagulation can potentially affect SOA mass yields in this study, both the 389 measured and coagulation-corrected size-dependent particle wall deposition coefficients 390 are used to correct for particle wall deposition in the  $\alpha$ -pinene ozonolysis experiments.

Assuming that the effect of coagulation is small, the particle wall deposition corrected number concentration data provide a test of the appropriateness of the particle wall deposition correction. The corrected number concentration should level off at a constant value (i.e., the initial particle number concentration), assuming no significant coagulation, when particle wall deposition is properly accounted for since the walldeposited particle number distribution is added to the suspended particle number distribution during particle wall loss correction. Neglecting coagulation, we account for 398 particle wall deposition in nucleation and low AS experiments using deposition 399 coefficients measured from the low AS-seed only experiments, while particle deposition 400 in high AS experiments are accounted for using coefficients measured from the high AS-401 seed only experiments. Figures S1 and S2 show the particle wall deposition-corrected 402 aerosol number and volume concentrations. Over all experiments, the particle wall 403 deposition-corrected final particle number concentration (i.e., at the end of the reaction) 404 is 9 to 17 % less than the initial particle number concentration for the low AS and high 405 AS experiments (Table S1), respectively, indicating that the particle wall deposition-406 corrected volume concentrations are slightly underestimated. The fact that the particle 407 wall deposition-corrected final particle number concentrations are somewhat smaller than 408 the initial particle number concentrations may be due to variations in particle wall 409 deposition rates in the AS-seed only and  $\alpha$ -pinene ozonolysis experiments or to 410 coagulation. To first examine variations in particle wall deposition rates, we used the 411 average of the measured low AS-seed only and high AS-seed only particle wall 412 deposition coefficients to account for particle wall deposition in all the experiments (Figs. 413 S3 and S4). While there is a negligible difference in the particle wall deposition corrected 414 volume concentrations (Figs. S3 and S4 vs. Figs. S1 and S2), a larger spread (1 to 22 %) 415 exists in the difference between the initial and final particle number concentrations when 416 the average measured particle wall deposition coefficients are used (Table S1). Therefore, 417 all subsequent nucleation and low AS data presented here are particle wall deposition-418 corrected using coefficients measured from the low AS-seed only experiments, and all 419 high AS data are corrected using particle wall deposition coefficients measured from the 420 high AS-seed only experiments. We furthermore conclude that variations in particle wall 421 deposition rates do not cause the decrease in the particle wall deposition-corrected final 422 number concentration and is most likely due to coagulation. Thus, the SOA data are also 423 corrected using the coagulation-corrected particle wall deposition coefficients (discussed 424 below). We show subsequently the relatively minor difference that correcting for 425 coagulation has on overall SOA mass yields. Therefore, we use SOA concentrations 426 corrected using the measured particle wall deposition coefficients for the bulk of the 427 analysis in this study.

428 Figure 2 shows the reaction profiles of the  $\alpha$ -pinene ozonolysis experiments. SOA 429 growth typically starts within 10 to 20 min of the start of the reaction. At either  $O_3$ 430 concentration, the molar ratio of  $O_3$  reacted to  $\alpha$ -pinene reacted is approximately 1:1 (i.e., 431 50 ppb  $\alpha$ -pinene reacted with 50 ppb O<sub>3</sub>), which indicates that O<sub>3</sub> reacts only with  $\alpha$ -432 pinene and not its oxidation products. As anticipated, the  $\alpha$ -pinene oxidation rates in the 433 100 ppb  $O_3$  experiments are significantly slower than those in the 500 ppb  $O_3$ 434 experiments. Figures 2a-c show that peak SOA levels are typically reached at reaction 435 time ~300 to 350 min in the 100 ppb  $O_3$  experiments, during which  $\ge$  95 % of the injected  $\alpha$ -pinene has reacted. In contrast, all the  $\alpha$ -pinene reacts within 80 to 90 min of 436 437 the start of reaction in the 500 ppb  $O_3$  experiments, and peak SOA levels are achieved at 438 reaction time  $\sim 100$  min (Figs. 2d-f). These results indicate that the O<sub>3</sub> concentration 439 dictates both the rate of  $\alpha$ -pinene oxidation and the time it takes to achieve peak SOA 440 growth.

441 Figure 3 shows the time-dependent growth curves (SOA mass concentration vs. 442  $\alpha$ -pinene reacted (Ng et al., 2006)) for the 100 and 500 ppb O<sub>3</sub> experiments. Only SOA growth data up to SOA peak concentrations are shown. SOA growth essentially stops 443 444 once all the  $\alpha$ -pinene has reacted. This is expected, as  $\alpha$ -pinene has only one double 445 bond; the first step of  $\alpha$ -pinene ozonolysis is rate-limiting and the first-generation 446 products are condensable (Ng et al., 2006; Chan et al., 2007). The time-dependent SOA 447 growth curves for experiments corresponding to different seed aerosol concentrations 448 overlap for both low and high  $O_3$  concentrations. This indicates that the initial AS seed 449 surface area does not influence the SOA growth rate within the range of AS seed surface 450 area concentration used. It is important to note that while it appears that the SOA growth 451 rate is faster in the 100 ppb O<sub>3</sub> relative to the 500 ppb O<sub>3</sub> experiments based on the time-452 dependent growth curves shown in Fig. 3, this is not the case. Instead, the observed time-453 dependent growth curves can be explained by the higher concentration of  $\alpha$ -pinene 454 having reacted during the 10 to 20 min delay of SOA formation in the 500 ppb  $O_3$ 455 experiments compared to the 100 ppb O<sub>3</sub> experiments (Fig. 2).

Figure 4 shows the time-dependent SOA mass yields as a function of initial total AS seed surface area for the 100 and 500 ppb O<sub>3</sub> experiments. Regardless of the O<sub>3</sub> 458 concentration, the SOA mass yields stay roughly constant despite the increase in AS seed 459 surface area. This indicates that the surface area concentration of AS seed aerosol does 460 not noticeably influence the partitioning of gas-phase  $\alpha$ -pinene ozonolysis products to the 461 particle phase within the range of AS seed surface area concentration used. Higher SOA 462 mass yields are observed in the 500 ppb  $O_3$  experiments, which indicates that the  $\alpha$ -463 pinene oxidation rate controls the absolute amount of SOA formed. It is important to note 464 that these conclusions are robust even when the average of the measured low AS-seed 465 only and high AS-seed only particle wall loss coefficients are used to account for particle 466 wall loss in all the experiments (Fig. S6). The enhancement of SOA mass yields at higher O<sub>3</sub> concentrations and the lack of a SOA mass yield dependence on AS seed surface area 467 468 (within the range of AS seed surface area concentration used in this study) will be 469 discussed further in Section 5.

The  $\alpha$ -pinene ozonolysis SOA mass yields obtained in this study are compared to those reported in previous studies in Fig. 5. Table S3 lists the experimental conditions employed in these studies. To facilitate comparison between the different studies, all the SOA mass yield and concentration data (including this study) are adjusted to an organic density of 1.0 g cm<sup>-3</sup>. As shown in Fig. 5, the SOA mass yields obtained at peak SOA growth in this study are generally consistent with those of previous studies where the chamber was operated in batch mode (that in this study).

477 To investigate the influence of coagulation on the SOA mass yields, the 478 coagulation-corrected size-dependent particle wall deposition coefficients are also used to 479 correct for particle wall deposition in the  $\alpha$ -pinene ozonolysis experiments. Specifically, 480 all nucleation and low AS data are particle wall deposition-corrected using coagulation-481 corrected coefficients derived from the low AS-seed only experiments, and all high AS 482 data are corrected using coagulation-corrected particle wall deposition coefficients 483 derived from the high AS-seed only experiments. Figure S7 shows the time-dependent 484 SOA mass yields (obtained using the coagulation-corrected and measured particle wall deposition coefficients) as a function of initial total AS seed surface area. SOA mass 485 486 yields obtained using the coagulation-corrected particle wall deposition coefficients are < 2 % (absolute values) higher than those using the measured particle wall deposition 487

488 coefficients. Similar to the SOA mass yields obtained using the measured particle wall 489 deposition coefficients (Figs. 4, S7c and S7d), SOA mass yields obtained using the 490 coagulation-corrected particle wall deposition coefficients stay roughly constant despite 491 the increase in AS seed surface area for both O<sub>3</sub> concentrations, and the SOA mass yields 492 are higher in the 500 ppb O<sub>3</sub> experiments (Figs. S7a and S7b). The mass yields obtained 493 at peak SOA growth are also generally consistent with those of previous studies (Fig. S8). 494 Taken together, this suggests that the effect of coagulation on the SOA mass yields is 495 likely minor for the aerosol concentrations used in this study. Therefore, only data that 496 have been particle wall deposition-corrected using coefficients measured in the low AS-497 seed only and high AS-seed only experiments are fitted to determine model parameters 498 for the vapor-particle dynamics model described in Section 3.

499 As noted earlier, optimal model values for  $\alpha_p$ ,  $\alpha_w$ ,  $\tau_{olig}$  and the branching ratios 500 between the oxidation products were determined by calculating the error between the 501 observed and modeled time-dependent SOA concentrations for all possible combinations 502 of model parameters. The combination of parameters with the lowest percent error is  $\alpha_w =$  $10^{-6}$ ,  $\alpha_p = 0.1$ ,  $\tau_{olig} = 4$  h, branching ratios = 0.6, 0.3, 0.05, 0.05 and 0 for oxidation 503 products with vapor pressures  $>10^3$ ,  $10^2$ , 10, 1 and  $0.1 \ \mu g \ m^{-3}$ , respectively. This 504 combination of parameters results in a percent error of 21% (Table S4). It is important to 505 506 note that predictions using  $\alpha_p = 0.1$  or 1 resulted in very similar errors; with the same combination of parameters and  $\alpha_p = 1$ , the percent error only increased to 22%. The "best 507 508 estimate" parameters determined following the Karnezi et al. (2014) method are as follows:  $\alpha_w = 3.6 \times 10^{-6}$ ,  $\alpha_p = 0.35$ ,  $\tau_{olig} = 6$  h, and branching ratios = 0.66, 0.16, 0.06, 0.06, 509 and 0.06 for oxidation products with vapor pressures  $>10^3$ ,  $10^2$ , 10, 1 and 0.1 µg m<sup>-3</sup>, 510 respectively. This combination of parameters results in an error of 37% (Table S4). 511 512 Model predictions using both sets of parameters are compared to measured SOA 513 concentrations in Fig. S9. The lowest-error parameters are used for the analysis in the remainder of this study. The best-fit  $\alpha_w = 10^{-6}$  corresponds to a first-order vapor-wall 514 deposition rate constant ( $k_{wall,on}$ ) of 10<sup>-4</sup> s<sup>-1</sup>. A wide range of vapor wall loss rates has 515 been reported (Figure 3 of Krechmer et al. (2016)). This  $k_{wall.on}$  value is comparable to 516 that reported by Matsunaga and Ziemann (2010) for a 8.2 m<sup>3</sup> chamber but significantly 517 518 faster than wall loss rates that have been measured in the Caltech chamber (Zhang et al.,

519 2015). The reason for this wide range of reported vapor wall loss rates is currently520 uncertain and outside the scope of this study.

#### 521 **5. Discussion**

522  $\alpha$ -pinene ozonolysis has been carried out at two O<sub>3</sub> mixing ratios (100 and 500 523 ppb) under varying AS seed aerosol surface area concentrations (0, ~1000 and ~3000  $\mu$ m<sup>2</sup> 524 cm<sup>-3</sup>).

## 525 5.1 Seed aerosol surface area effect

526 Figure 3 shows that the time-dependent SOA growth curves for experiments with 527 different seed area concentrations overlap at both  $O_3$  concentrations, which indicates the 528 AS seed surface area does not affect the rate of SOA growth within the range of AS seed 529 surface area concentration used in this study. This observation differs from findings by 530 Pathak et al. (2007b) for the  $O_3+\alpha$ -pinene system, who showed that even though the final 531 SOA mass yields measured in the reaction of 7.3 ppb  $\alpha$ -pinene with 1500 ppb O<sub>3</sub> were 532 similar in their seeded and unseeded experiments, SOA growth was considerably slower 533 in unseeded experiments compared to seeded experiments. The authors suggested that the 534 slow SOA formation rate in their unseeded experiment was the result of SOA formation 535 being limited by the mass transfer of semi-volatile oxidation products to newly formed 536 particles (via nucleation) during the early stages of the experiment. These newly formed 537 particles have a significantly smaller aerosol surface area for gas-particle partitioning as 538 compared to that of seed aerosol particles in the seeded experiments. Consequently, the 539 semi-volatile oxidation products accumulated in the gas phase during the early stages of 540 the unseeded experiments, resulting in slower SOA growth compared to the seeded 541 experiments. The observation that the presence of seed aerosol does not influence the 542 SOA growth rate in the present study may be explained by the relatively high 543 concentrations of a-pinene reacted and SOA mass loadings obtained. Previous studies 544 have shown that the delay between the onset of VOC oxidation and SOA formation in 545 unseeded experiments is most pronounced at low aerosol loadings (Kroll et al., 2007). 546 We note that the concentrations of  $\alpha$ -pinene reacted and SOA mass loadings obtained in 547 this study are significantly larger than those reported by Pathak et al. (2007b). Therefore, 548 it is possible that due to the relatively large concentrations of  $\alpha$ -pinene reacted in this 549 study, substantial concentrations of gas-phase oxidation products are generated, which 550 results in rapid partitioning into the particle phase even in the absence of seed aerosol. 551 This is evident from the large increase in the particle number concentration during the 552 early stages of the unseeded 100 and 500 ppb O<sub>3</sub> experiments, where the particle number concentration increased to ~8000 and ~10000 particles/cm<sup>3</sup> during the first 45 min of the 553 554 100 and 500 ppb O<sub>3</sub> experiments, respectively (Fig. S1a and S2a). Thus, the SOA growth 555 rates are not controlled by the presence of AS seed in this study.

Figure 4 shows that for both  $O_3$  mixing ratios used, the time-dependent SOA mass yield is similar at any given AS seed surface area (see also Table 1). The absence of a SOA growth dependence on the AS seed surface area is similar to observations reported by McVay et al. (2016) for the  $\alpha$ -pinene photooxidation (OH-driven chemistry) system, but differ from those reported by Zhang et al. (2014) for the toluene photooxidation system in which the SOA mass yield increased with the surface area concentration of seed aerosol.

563 The best-fit  $\alpha_p = 0.1$  (or  $\alpha_p = 1$ , with almost the same percentage error) suggests 564 the absence of significant limitations to vapor-particle mass transfer in the present  $\alpha$ -565 pinene ozonolysis study, and that SOA formation is governed by quasi-equilibrium 566 growth (Saleh et al., 2013; McVay et al., 2014), which occurs when SOA-forming vapors 567 are produced at a rate that is significantly slower than that required to establish gas-568 particle equilibrium (Shiraiwa and Seinfeld, 2012; Zhang et al., 2012). Moreover, the 569 characteristic timescale to establish gas-particle equilibrium is less than those for reaction 570 and vapor-wall deposition. When the vapor and particle phases maintain equilibrium, 571 gas-particle equilibrium is controlled by the amount of organic matter in the VOC 572 system. As a result, the rate of condensation of SOA-forming vapors is independent of 573 the seed aerosol surface area (McVay et al., 2014). The best-fit  $\alpha_p = 0.1$  is within the 574 range of  $\alpha_p$  coefficients determined from  $\alpha$ -pinene ozonolysis SOA thermodenuder 575 studies ( $\alpha_p = 0.1$ ) (Saleh et al., 2013; Saha et al., 2016) and  $\alpha$ -pinene photooxidation 576 chamber studies ( $\alpha_p = 0.1$  or 1) (McVay et al., 2016). Notably, this result differs markedly 577 from that for toluene photooxidation (Zhang et al., 2014), where  $\alpha_p$  was determined to be 578 0.001, and for which, since the SOA mass yield was strongly dependent on the seed 579 aerosol surface area, the condensation of SOA-forming vapors onto seed aerosol particles 580 was kinetically limited (McVay et al., 2014). Kinetically-limited SOA growth occurs 581 when the timescale for gas-particle equilibrium is competitive with or exceeds the 582 timescale for reaction and vapor wall deposition, and may reflect imperfect 583 accommodation of gas-phase organics to the particle phase. The markedly different 584 behavior of the  $\alpha$ -pinene and toluene SOA systems could be due to differences in SOA 585 volatility and aerosol physical phase state (McVay et al., 2016).

### 586 **5.2 Oxidation rate effect**

587 At higher  $O_3$  concentrations, the  $\alpha$ -pinene oxidation rate increases, leading to 588 higher SOA mass yields (the "oxidation rate effect"). This behavior was previously 589 observed by Ng et al. (2007) for the *m*-xylene photooxidation system, for which the 590 oxidation rate effect was attributed to the loss of semi-volatile condensable products to 591 chamber walls in competition with condensation onto seed particles to form SOA.

592 SOA formation from  $\alpha$ -pinene ozonolysis is presumed to be driven by a range of 593 semi- and low-volatility first-generation products arising from reaction of O<sub>3</sub> with the 594 single C=C double bond (Ng et al., 2006). These products are subject to two competing 595 routes: condensation to particles to form SOA or deposition on the chamber walls. Each 596 process can be represented in terms of a first-order rate constant:  $k_{wall,on}$  and  $k_{particle,on}$  (s<sup>-1</sup>). The rate of vapor-wall deposition of condensable species A is then  $k_{wall on} \times [A]$  (molec 597 cm<sup>-3</sup> s<sup>-1</sup>) and the rate of condensation onto particles is  $k_{particle,on} \times [A]$  (molec cm<sup>-3</sup> s<sup>-1</sup>). 598 599 Increasing the rate of reaction increases the concentration of [A], but the relative rates of 600 vapor-wall deposition and condensation onto particles will remain the same. In general, 601 however, both vapor-wall deposition and vapor-particle condensation are reversible 602 processes (McVay et al., 2014; Zhang et al., 2014). The first-order rate constant for 603 evaporation from the wall can be represented as (Matsunaga and Ziemann, 2010):

$$k_{wall,off} = k_{wall,on} \left(\frac{C_i^*}{C_w}\right)$$
(15)

21

605 where  $C_i^*$  is the saturation concentration and  $C_w$  is the assumed equivalent wall organic 606 concentration. The rate of evaporation from particles is:

$$k_{particle,off} = k_{particle,on} \left(\frac{C_i^*}{C_{aer}}\right)$$
(16)

608 where  $C_{aer}$  is the organic aerosol concentration ( $C_{aer}=\sum A_k+M_{init}$ ).

609 The difference between  $C_{aer}$  and  $C_w$  is the key to explaining the oxidation rate effect. At the beginning of the experiment, Caer is very small because the inorganic seeds 610 are essentially non-absorbing. Therefore,  $k_{particle,off}$  is large, and the net SOA growth is 611 small. In contrast,  $C_w$  is considered to be substantial (on the order of 10 mg m<sup>-3</sup>) and to be 612 613 essentially constant throughout the experiment (Matsunaga and Ziemann, 2010; McVay 614 et al., 2014; Zhang et al., 2014). Model predictions are insensitive to the value of  $C_w$ since, in any event,  $C_w$  is significantly larger than  $C_{aer}$  (Zhang et al., 2014). Therefore, 615 616  $k_{wall.off}$  is small at the beginning of the experiment and the net vapor wall loss rate is fast. 617 As  $C_{aer}$  increases, the net SOA condensation rate increases relative to the net vapor wall 618 loss rate. When the reaction rate increases corresponding to higher  $O_3$  concentrations,  $C_{aer}$ 619 grows more quickly because more condensable species are available to form SOA, and 620 the net condensation rate increases more rapidly. Therefore, the observed oxidation rate 621 effect is due to vapor wall deposition, and arises because vapor-particle and vapor-wall 622 condensation are essentially reversible processes. This explanation is consistent with 623 simulations varying the O<sub>3</sub> concentration in which all species are non-volatile (i.e., do not 624 evaporate from the particles or the wall). In this case, no oxidation rate effect is observed as the  $O_3$  concentration increases. The growth curves for different  $O_3$  concentrations 625 626 overlap, and the same yield is obtained regardless of O<sub>3</sub> concentration (Fig. S10).

Sensitivity tests were performed to determine the point at which SOA formation is no longer influenced by the O<sub>3</sub> concentration. In these simulations, the initial  $\alpha$ -pinene concentration is fixed at 48 ppb, while the O<sub>3</sub> concentration is varied from 75 to 1000 ppb. The rate of O<sub>3</sub> injection is assumed to remain constant as the O<sub>3</sub> concentration is increased to mimic the experimental protocol (i.e., O<sub>3</sub> injection time is increased to achieve higher O<sub>3</sub> concentrations). The O<sub>3</sub> injection rate used in these simulations is fixed

at 500/54.25 ppb min<sup>-1</sup>, which is the same as that used to analyze results from the 500 633 634 ppb  $O_3$  experiments. Model predictions in Fig. S11 show that the maximum SOA mass 635 concentration increases with increasing O<sub>3</sub> concentration up to approximately 500 ppb O<sub>3</sub>. Beyond this O<sub>3</sub> concentration, the SOA growth curves overlap and the maximum 636 637 SOA mass concentration does not increase even when more  $O_3$  is added. This plateau 638 arises due to the lengthening time required to inject increasing amounts of O<sub>3</sub>. More than 639 1 h is required to inject > 500 ppb of  $O_3$ , and by this time, virtually all of the  $\alpha$ -pinene has 640 reacted. Increasing the  $O_3$  concentration after all of the  $\alpha$ -pinene has reacted does not lead 641 to any changes in the SOA mass concentration. However, if a faster injection rate of  $O_3$  is used, the oxidation rate effect will persist to higher  $O_3$  concentrations (i.e., > 500 ppb  $O_3$ ) 642 643 (Fig S12). With a faster injection rate, 500 ppb  $O_3$  is injected before all of the  $\alpha$ -pinene 644 has reacted. Continuing to inject  $O_3$  to a higher concentration (i.e., 750 ppb) will cause  $\alpha$ -645 pinene to decay faster and SOA to grow faster than when the  $O_3$  injection stops at 500 646 ppb. The oxidation rate effect is then apparent at higher  $O_3$  concentrations. If, instead of 647 using an injection rate of  $O_3$ , simulations are run using fixed initial  $O_3$  (not possible 648 experimentally), the rate effect persists to even higher O<sub>3</sub> concentrations. The relative 649 increase in yield with increasing O<sub>3</sub> concentrations slows at very high O<sub>3</sub> concentrations 650 because the rate of reaction becomes substantially faster than the vapor wall deposition 651 rate, and there is less marginal effect to increasing the reaction rate.

652 It should be noted that while we showed that the observed oxidation rate effect 653 (i.e., higher SOA mass yields as a result of faster hydrocarbon oxidation rates) is a 654 consequence of vapor-wall deposition, the possibility that differing peroxy radical (RO<sub>2</sub>) 655 chemistry in the 100 and 500 ppb O<sub>3</sub> experiments may play some role in influencing the 656 SOA mass yields cannot be discounted.  $RO_2$  radicals, which are formed from the 657 decomposition of excited Criegee intermediates (Docherty et al., 2005), may be produced 658 at faster rates in the 500 ppb  $O_3$  experiments. This may lead to the higher production of 659 condensable oxidation products from the  $RO_2+RO_2$  reaction pathway in the 500 ppb  $O_3$ 660 experiments (relative to those formed in the 100 ppb O<sub>3</sub> experiments), which may result 661 in higher SOA mass yields.

# 662 5.3 Interplay of the seed aerosol surface area effect and the oxidation rate effect

663 In this study, we observe an oxidation rate effect but not a seed aerosol surface 664 area effect. In Zhang et al. (2014), a seed aerosol surface area effect was observed, but 665 the variation of the oxidation rate was not studied. A key aspect of vapor wall deposition is the potential interplay between the seed aerosol surface area effect and the oxidation 666 667 rate effect. To examine this interplay in the  $\alpha$ -pinene ozonolysis system, simulations were 668 carried out by varying the seed aerosol surface area and the O<sub>3</sub> concentration 669 simultaneously, while using the branching ratios, oligomerization rate, and vapor wall 670 deposition rate parameters obtained in the present study. The initial  $\alpha$ -pinene 671 concentration was set to 50 ppb, and a fixed O<sub>3</sub> concentration was used in place of a 672 linear injection.  $\alpha_p$  was varied at 0.001, 0.01, 0.1, and 1 in these simulations. Figure 6 673 shows the SOA mass yield at peak SOA growth as a function of both the seed aerosol surface area and O<sub>3</sub> concentration for  $\alpha_p = 1, 0.1, 0.01$ , and 0.001. For  $\alpha_p = 1$  or 0.1, the 674 675 oxidation rate dominates: SOA mass yield increases significantly as O<sub>3</sub> concentration 676 increases while the seed aerosol surface area has a negligible effect. For  $\alpha_p = 0.01$ , both effects can be observed in different regions: at low O3 concentrations and high seed 677 678 aerosol surface areas, the oxidation rate effect dominates; at low seed aerosol surface 679 areas and high O<sub>3</sub> concentrations, the seed surface area dominates. At low seed aerosol surface areas and low O<sub>3</sub> concentrations, both effects are present. For  $\alpha_p = 0.001$ , the seed 680 681 aerosol surface area effect dominates except at very high seed aerosol surface areas. 682 These observations show that the presence of an oxidation rate effect and/or seed aerosol surface area effect depends on a complex interplay of factors, such as  $\alpha_p$ , the rate of 683 684 hydrocarbon oxidation, and the amount of seed surface area present.

## 685 **6. Implications**

In this study, we systematically examine the roles of gas-particle partitioning and VOC oxidation rate in the presence of vapor-wall deposition in  $\alpha$ -pinene ozonolysis. We show that despite the presence of vapor-wall deposition, SOA mass yields at peak SOA growth remain approximately constant regardless of the seed aerosol surface area (within the range of AS seed surface area concentration used in this study). This observation is consistent with SOA formation in the  $\alpha$ -pinene ozonolysis system being governed by quasi-equilibrium growth, for which there are no substantial limitations to vapor-particle 693 mass transfer. This result was demonstrated in a previous modeling study which showed 694 that increasing the seed-to-chamber surface area ratio will lead to increased SOA growth 695 only in cases in which the condensation of SOA-forming vapors onto seed aerosol 696 particles is kinetically limited as a result of imperfect accommodation of gas-phase 697 organics to the particle phase (McVay et al., 2014).

An important implication of this study is that diverting vapor-wall deposition in chamber studies via the addition of ever-increasing quantities of seed aerosol particles is not effective in VOC systems for which SOA formation is governed by quasi-equilibrium growth. This study also underscores the importance of accounting for particle wall deposition appropriately in chamber studies, to avoid erroneous conclusions regarding the role of gas-particle partitioning (quasi-equilibrium vs. kinetically-limited SOA growth) in influencing vapor wall loss in the VOC system.

705 We note that the present study shows that the SOA mass yield is independent of seed aerosol surface area concentration for values ranging from 0 to  $\sim 3000 \ \mu m^2 \ cm^{-3}$ . 706 This corresponds to a seed-to-chamber surface area ratio of 0 to  $\sim 1 \times 10^{-3}$ , which is 707 708 substantially smaller than the range used by Zhang et al. (2014) to study the influence of 709 vapor-wall deposition on toluene photooxidation SOA formation in the Caltech chamber (i.e., 0 to  $\sim 5 \times 10^{-3}$ ). It is possible that a SOA mass yield dependence on the seed surface 710 711 area may have become more apparent had a larger range of seed aerosol surface area (i.e.,  $> 3000 \text{ }\mu\text{m}^2 \text{ cm}^{-3}$ ), and hence a larger range of seed-to-chamber surface area ratio, been 712 713 used here. One consideration is that coagulation may become increasingly important, and 714 will need to be accounted for, when higher seed aerosol number concentrations (relative 715 to those used in this study) are used (Seinfeld and Pandis, 2006; Pierce et al., 2008). A detailed analysis of the effect of seed aerosol surface area concentrations  $> 3000 \text{ }\mu\text{m}^2 \text{ cm}^-$ 716 <sup>3</sup> on  $\alpha$ -pinene ozonolysis SOA mass yields will be the subject of forthcoming work. 717

Higher SOA mass yields at peak SOA growth are observed in the present study when  $O_3$  is increased from 100 to 500 ppb. This is because  $\alpha$ -pinene is oxidized more quickly, which leads to gas-phase oxidation products being formed more rapidly, and consequently partitioning more quickly onto AS seed aerosol particles before they are lost to the chamber walls. Therefore, the oxidation rate effect (i.e., higher SOA mass 723 yields as a result of faster hydrocarbon oxidation rates) is a consequence of vapor-wall 724 deposition. An important implication of this study is that SOA mass yields can be 725 affected by vapor-wall deposition in VOC systems that are not characterized by slow 726 mass accommodation of gas-phase organics to the particle phase (Zhang et al., 2014). 727 Thus, this work demonstrates that the effect of vapor-wall deposition on SOA mass yields 728 can be mitigated through the use of excess oxidant concentrations. It should be noted that 729 the  $\alpha$ -pinene ozonolysis SOA mass yields (absolute values) increased by 5 to 9 % when  $O_3$  is increased from 100 to 500 ppb (for an initial  $\alpha$ -pinene concentration of ~50 ppb), 730 731 where SOA formation is governed by quasi-equilibrium growth. In the absence of vapor-732 wall deposition, SOA mass yields are predicted by the model used here to approximately 733 double from those observed experimentally. In contrast, Zhang et al. (2014) showed that 734 the presence of vapor-wall deposition led to underestimation of SOA formation by factors 735 as much as four in the toluene photooxidation system, where the condensation of SOA-736 forming vapors onto seed aerosol is kinetically limited. Taken together, these results 737 indicate that the magnitude by which vapor-wall deposition affects SOA mass yields 738 depends on the extent to which the VOC system is governed by kinetically-limited SOA 739 condensational growth.

740 Given these observations of how gas-particle partitioning can influence the 741 magnitude by which vapor-wall deposition affects SOA mass yields, an overriding 742 question is: what controls the gas-particle partitioning behavior of SOA formed in 743 different VOC systems?  $\alpha_p$  describes the overall mass transfer of vapor molecules into the 744 particle phase (McVay et al., 2014; Zhang et al., 2014). Thus,  $\alpha_p$  affects the vapor-745 particle equilibrium timescale, which, depending on the extent to which it is competitive 746 with the timescales for reaction and vapor-wall deposition, determines whether SOA 747 formation is governed by kinetically-limited or quasi-equilibrium growth. Markedly 748 different  $\alpha_p$  values could arise from the physical phase state of the SOA formed. As 749 discussed by McVay et al. (2014, 2016), if the SOA formed exists in a semi-solid state 750 (Vaden et al., 2010; Virtanen et al., 2010; Cappa and Wilson, 2011; Vaden et al., 2011; 751 Virtanen et al., 2011; Kuwata and Martin, 2012; Perraud et al., 2012; Saukko et al., 2012; Abramson et al., 2013; Renbaum-Wolff et al., 2013), a low value of  $\alpha_p$  might be expected 752 753 owing to retarded surface accommodation and particle-phase diffusion (Zaveri et al.,

2014). Quantification of  $\alpha_p$  is challenging experimentally, and reported  $\alpha_p$  values for the same system can vary by several orders of magnitude (Grieshop et al., 2007; Stanier et al., 2007; Vaden et al., 2011; Miles et al., 2012; Saleh et al., 2013 Saha et al., 2016). Therefore,  $\alpha_p$  of SOA formed in different VOC systems need to be better constrained through a combination of experimental and modeling efforts.

759 While not investigated in detail in this study, the timescale of oligomerization may play an important role depending on the SOA growth regime (i.e., kinetically-760 761 limited vs. quasi-equilibrium). Currently, it is unclear how the timescale of 762 oligomerization in a VOC system where SOA formation is dominated by quasi-763 equilibrium growth (e.g.  $\alpha$ -pinene ozonolysis) may differ from one that is dominated by 764 kinetically-limited growth (e.g. toluene photooxidation (Zhang et al., 2014)). This 765 requires further investigation through a combination of experimental and modeling efforts to improve our understanding of how particle-phase processes (e.g. 766 767 oligomerization) affect gas-particle partitioning, and consequently influence the 768 magnitude by which vapor-wall deposition affects SOA mass yields.

769 The SOA mass yield from the ozonolysis of monoterpenes in the GEOS-CHEM chemical transport model (19 % at 10  $\mu$ g m<sup>-3</sup>) is currently based on that measured in  $\alpha$ -770 771 pinene ozonolysis studies by Shilling et al. (2008) (Pye et al., 2010). Shilling et al. (2008) 772 measured these SOA mass yields in a teflon chamber operated in continuous-flow mode, 773 as opposed to batch mode, which is how experiments in the present study and most of 774 those shown in Fig. 5 and Table S3 were conducted. While it is not possible to directly 775 compare our results with those of Shilling et al. (2008) due to differences in SOA mass 776 concentrations, the SOA mass concentrations and yields measured in the current study 777 are generally consistent with those of previous batch chamber studies. The SOA mass vields at  $\sim 10 \text{ µg m}^{-3}$  SOA mass concentration measured by Shilling et al. (2008) are 778 779 generally higher than those measured in chambers operated in batch mode (Griffin et al., 780 1999; Cocker et al., 2001b; Presto et al., 2005; Presto and Donahue, 2006; Pathak et al., 781 2007b) (Fig. 5). One possible explanation for the higher SOA mass yields in the 782 continuous-flow, steady state, mode is that the SOA-forming vapors are in equilibrium 783 with the chamber walls and seed aerosol, hence minimizing the irreversible loss of SOA-

784 forming vapors to the chamber walls (Shilling et al., 2008). However, the extent to which 785 SOA mass yields obtained in a continuous-flow reactor are influenced by vapor wall loss 786 is unclear. Using a continuous-flow reactor, Ehn et al. (2014) observed  $\alpha$ -pinene 787 ozonolysis SOA mass yields to increase with increasing seed aerosol surface area but 788 required  $\alpha_p = 1$  to fit the observed SOA growth. The observed vapor-wall deposition rate constant in their continuous-flow reactor  $(0.011 \text{ s}^{-1})$  is two orders of magnitude larger 789 than that of the GTEC chamber  $(10^{-4} \text{ s}^{-1})$ . The estimated timescales for gas-particle and 790 791 gas-wall partitioning are also approximately equal in their continuous-flow reactor. This 792 indicates that SOA condensational growth is kinetically limited in their continuous-flow 793 reactor even at  $\alpha_p = 1$  (Ehn et al., 2014; McVay et al., 2014), which suggests that SOA 794 mass yields measured in their continuous-flow reactor may be significantly affected by 795 vapor-wall deposition.

796 Previous studies on SOA formation from the OH and NO<sub>3</sub> oxidation of biogenic 797 VOCs have similarly reported higher SOA mass yields in the presence of higher oxidant 798 concentrations. For example, in the NO<sub>3</sub> oxidation of  $\beta$ -pinene, Boyd et al. (2015) 799 reported SOA mass yields 10 to 30 % higher than those previously reported by Fry et al. 800 (2009, 2014). In addition to differences in the experimental conditions of the two studies 801 (which may lead to differing RO<sub>2</sub> chemistry), Boyd et al. (2015) hypothesized that the 802 higher SOA mass yields could also be a result of the higher NO<sub>3</sub> concentrations used in 803 their study (which led to faster  $\beta$ -pinene oxidation rates) compared to those used by Fry 804 et al. (2009, 2014). The oxidation rate effect was also observed in the *m*-xylene 805 photooxidation system, where Ng et al. (2007) showed that the SOA mass yields were 806 dependent on the *m*-xylene oxidation rate, with higher OH concentrations (and hence 807 faster oxidation rates) resulting in higher SOA mass yields. The authors dismissed the 808 possibility of the different SOA mass yields being a result of different RO<sub>2</sub> chemistry 809 since all their *m*-xylene photooxidation experiments were performed under high-NO<sub>x</sub> 810 conditions and the RO<sub>2</sub> reacted virtually entirely with NO. Together, these studies show 811 that faster hydrocarbon oxidation rates can alleviate the effects of vapor-wall deposition 812 on SOA mass yields in different VOC systems.

813 This gives rise to the question: should chamber SOA experiments on different 814 VOC systems be performed under as rapid oxidation conditions as possible (i.e., large 815 oxidant concentrations) to reduce the effects of vapor-wall deposition? A recent study by 816 McVay et al. (2016) reported similar SOA growth under low and high OH levels for α-817 pinene photooxidation. The authors hypothesized that the autoxidation mechanism likely 818 becomes a more important pathway at low OH levels (Crounse et al., 2013), and thus 819 contributes substantially to SOA growth. Therefore, it is possible that certain reaction 820 pathways and mechanisms (which are important in the atmosphere) are biased when 821 unusually high levels of oxidants are used in chamber experiments (e.g. autoxidation). 822 Thus, this underscores the need to design chamber experiments that simultaneously 823 mitigate the magnitude of vapor-wall deposition while ensuring that reaction conditions, 824 and consequently reaction pathways and oxidation products, are atmospherically relevant. 825 More importantly, the impact of vapor-wall deposition on SOA formation and evolution 826 in various VOC systems conducted under different reaction conditions (regardless of 827 atmospheric relevance) needs to be quantified through a combination of experimental and 828 modeling efforts. Similar to this study, experiments should be performed using different 829 seed aerosol surface area and oxidant concentrations to study their influence on vapor-830 wall deposition and SOA mass yields. If the effects of vapor-wall loss are found to be 831 strongly dependent on seed aerosol surface area and/or oxidant concentrations (e.g. 832 toluene photooxidation, where SOA formation may be underestimated by factors as much 833 as four (Zhang et al., 2014)), further experiments aimed at measuring the wall deposition 834 rates of the oxidation products should be performed. These wall deposition rates can then 835 be used in predictive models to determine the vapor-wall and vapor-particle mass 836 accommodation coefficients of these oxidation products. Consequently, this will allow us 837 to determine the fraction of SOA-forming vapors partitioning to the particle phase vs. lost 838 to the chamber walls (Zhang et al., 2015; Krechmer et al., 2016).

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## 847 **References**

Abramson, E., Imre, D., Beranek, J., Wilson, J., and Zelenyuk, A.: Experimental
determination of chemical diffusion within secondary organic aerosol particles, Phys.
Chem. Chem. Phys., 15, 2983-2991, 10.1039/c2cp44013j, 2013.

851 Bahreini, R., Keywood, M. D., Ng, N. L., Varutbangkul, V., Gao, S., Flagan, R. C.,

852 Seinfeld, J. H., Worsnop, D. R., and Jimenez, J. L.: Measurements of Secondary Organic

853 Aerosol from Oxidation of Cycloalkenes, Terpenes, and m-Xylene Using an Aerodyne

- Aerosol Mass Spectrometer, Environmental Science & Technology, 39, 5674-5688,
- 855 10.1021/es048061a, 2005.
- Boyd, C. M., Sanchez, J., Xu, L., Eugene, A. J., Nah, T., Tuet, W. Y., Guzman, M. I., and
  Ng, N. L.: Secondary organic aerosol formation from the β-pinene+NO3 system: effect of
  humidity and peroxy radical fate, Atmos. Chem. Phys., 15, 7497-7522, 10.5194/acp-157497-2015, 2015.
- Canagaratna, M. R., Jimenez, J. L., Kroll, J. H., Chen, Q., Kessler, S. H., Massoli, P.,
  Hildebrandt Ruiz, L., Fortner, E., Williams, L. R., Wilson, K. R., Surratt, J. D., Donahue,
  N. M., Jayne, J. T., and Worsnop, D. R.: Elemental ratio measurements of organic
  compounds using aerosol mass spectrometry: characterization, improved calibration, and
  implications, Atmos. Chem. Phys., 15, 253-272, 10.5194/acp-15-253-2015, 2015.
- 865 Capouet, M., Müller, J. F., Ceulemans, K., Compernolle, S., Vereecken, L., and Peeters,
- 866 J.: Modeling aerosol formation in alpha-pinene photo-oxidation experiments, Journal of
- 867 Geophysical Research: Atmospheres, 113, n/a-n/a, 10.1029/2007JD008995, 2008.
- 868 Cappa, C. D., and Wilson, K. R.: Evolution of organic aerosol mass spectra upon heating:
- 869 implications for OA phase and partitioning behavior, Atmospheric Chemistry and
  870 Physics, 11, 1895-1911, 10.5194/acp-11-1895-2011, 2011.

- 871 Cappa, C. D., Jathar, S. H., Kleeman, M. J., Docherty, K. S., Jimenez, J. L., Seinfeld, J.
- H., and Wexler, A. S.: Simulating secondary organic aerosol in a regional air quality
- 873 model using the statistical oxidation model Part 2: Assessing the influence of vapor
- wall losses, Atmos. Chem. Phys., 16, 3041-3059, 10.5194/acp-16-3041-2016, 2016.
- 875 Carlton, A. G., Bhave, P. V., Napelenok, S. L., Edney, E. D., Sarwar, G., Pinder, R. W.,
- 876 Pouliot, G. A., and Houyoux, M.: Model Representation of Secondary Organic Aerosol in
- 877 CMAQv4.7, Environmental Science & Technology, 44, 8553-8560, 10.1021/es100636q,
- 878 2010.
- Chan, A. W. H., Kroll, J. H., Ng, N. L., and Seinfeld, J. H.: Kinetic modeling of
  secondary organic aerosol formation: effects of particle- and gas-phase reactions of
  semivolatile products, Atmospheric Chemistry and Physics, 7, 4135-4147, 2007.
- Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: State-of-the-art chamber facility for
  studying atmospheric aerosol chemistry, Environmental Science & Technology, 35,
  2594-2601, 10.1021/es0019169, 2001a.
- 885 Cocker, D. R., Clegg, S. L., Flagan, R. C., and Seinfeld, J. H.: The effect of water on gas-
- 886 particle partitioning of secondary organic aerosol. Part I: alpha-pinene/ozone system,
- 887 Atmospheric Environment, 35, 6049-6072, 10.1016/s1352-2310(01)00404-6, 2001b.
- Crounse, J. D., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., and Wennberg, P. O.:
  Autoxidation of Organic Compounds in the Atmosphere, J. Phys. Chem. Lett., 4, 35133520, 10.1021/jz4019207, 2013.
- Krump, J. G., and Seinfeld, J. H.: Turbulent Deposition and Gravitational Sedimentation
  of an Aerosol in a Vessel of Arbitrary Shape, Journal of Aerosol Science, 12, 405-415,
  10.1016/0021-8502(81)90036-7, 1981.
- DeCarlo, P. F., Slowik, J. G., Worsnop, D. R., Davidovits, P., and Jimenez, J. L.: Particle
  morphology and density characterization by combined mobility and aerodynamic
  diameter measurements. Part 1: Theory, Aerosol Sci. Technol., 38, 1185-1205,
  10.1080/027868290903907, 2004.

- 898 DeCarlo, P. F., Kimmel, J. R., Trimborn, A., Northway, M. J., Jayne, J. T., Aiken, A. C.,
- 899 Gonin, M., Fuhrer, K., Horvath, T., Docherty, K. S., Worsnop, D. R., and Jimenez, J. L.:
- 900 Field-deployable, high-resolution, time-of-flight aerosol mass spectrometer, Analytical
- 901 Chemistry, 78, 8281-8289, 10.1021/ac061249n, 2006.
- 902 Docherty, K. S., Wu, W., Lim, Y. B., and Ziemann, P. J.: Contributions of organic
- 903 peroxides to secondary aerosol formed from reactions of monoterpenes with O3,
- 904 Environmental Science & Technology, 39, 4049-4059, 10.1021/es050228s, 2005.
- Donahue, N. M., Robinson, A. L., Stanier, C. O., and Pandis, S. N.: Coupled partitioning,
  dilution, and chemical aging of semivolatile organics, Environmental Science &
- 907 Technology, 40, 2635-2643, 10.1021/es052297c, 2006.
- 908 Donahue, N. M., Epstein, S. A., Pandis, S. N., and Robinson, A. L.: A two-dimensional
- 909 volatility basis set: 1. organic-aerosol mixing thermodynamics, Atmospheric Chemistry
- 910 and Physics, 11, 3303-3318, 10.5194/acp-11-3303-2011, 2011.
- 911 Ehn, M., Thornton, J. A., Kleist, E., Sipila, M., Junninen, H., Pullinen, I., Springer, M.,
- 912 Rubach, F., Tillmann, R., Lee, B., Lopez-Hilfiker, F., Andres, S., Acir, I. H., Rissanen,
- 913 M., Jokinen, T., Schobesberger, S., Kangasluoma, J., Kontkanen, J., Nieminen, T.,
- 914 Kurten, T., Nielsen, L. B., Jorgensen, S., Kjaergaard, H. G., Canagaratna, M., Dal Maso,
- 915 M., Berndt, T., Petaja, T., Wahner, A., Kerminen, V. M., Kulmala, M., Worsnop, D. R.,
- 916 Wildt, J., and Mentel, T. F.: A large source of low-volatility secondary organic aerosol,
- 917 Nature, 506, 476-479, 10.1038/nature13032, 2014.
- Fry, J. L., Kiendler-Scharr, A., Rollins, A. W., Wooldridge, P. J., Brown, S. S., Fuchs,
  H., Dube, W., Mensah, A., dal Maso, M., Tillmann, R., Dorn, H. P., Brauers, T., and
  Cohen, R. C.: Organic nitrate and secondary organic aerosol yield from NO3 oxidation of
  beta-pinene evaluated using a gas-phase kinetics/aerosol partitioning model, Atmospheric
- 922 Chemistry and Physics, 9, 1431-1449, 2009.
- 923 Fry, J. L., Draper, D. C., Barsanti, K. C., Smith, J. N., Ortega, J., Winkler, P. M., Lawler,
- 924 M. J., Brown, S. S., Edwards, P. M., Cohen, R. C., and Lee, L.: Secondary Organic
- 925 Aerosol Formation and Organic Nitrate Yield from NO3 Oxidation of Biogenic

- 926 Hydrocarbons, Environmental Science & Technology, 48, 11944-11953,
  927 10.1021/es502204x, 2014.
- Gao, S., Ng, N. L., Keywood, M., Varutbangkul, V., Bahreini, R., Nenes, A., He, J. W.,
  Yoo, K. Y., Beauchamp, J. L., Hodyss, R. P., Flagan, R. C., and Seinfeld, J. H.: Particle
  phase acidity and oligomer formation in secondary organic aerosol, Environmental
- 931 Science & Technology, 38, 6582-6589, 10.1021/es049125k, 2004.
- Grieshop, A. P., Donahue, N. M., and Robinson, A. L.: Is the gas-particle partitioning in
  alpha-pinene secondary organic aerosol reversible?, Geophys. Res. Lett., 34, n/a-n/a,
  10.1029/2007GL029987, 2007.
- 935 Griffin, R. J., Cocker, D. R., Flagan, R. C., and Seinfeld, J. H.: Organic aerosol formation
- from the oxidation of biogenic hydrocarbons, J. Geophys. Res.-Atmos., 104, 3555-3567,
  10.1029/1998jd100049, 1999.
- 938 Guenther, A. B., Jiang, X., Heald, C. L., Sakulyanontvittaya, T., Duhl, T., Emmons, L.
- 939 K., and Wang, X.: The Model of Emissions of Gases and Aerosols from Nature version
- 940 2.1 (MEGAN2.1): an extended and updated framework for modeling biogenic emissions,
- 941 Geoscientific Model Development, 5, 1471-1492, 10.5194/gmd-5-1471-2012, 2012.
- Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
  Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann,
  H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr,
  A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A., Prevot, A. S. H., Seinfeld,
  J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The formation, properties and impact
  of secondary organic aerosol: current and emerging issues, Atmospheric Chemistry and
  Physics, 9, 5155-5236, 2009.
- Henry, K. M., Lohaus, T., and Donahue, N. M.: Organic Aerosol Yields from alphaPinene Oxidation: Bridging the Gap between First-Generation Yields and Aging
  Chemistry, Environmental Science & Technology, 46, 12347-12354, 10.1021/es302060y,
  2012.

- Hildebrandt, L., Donahue, N. M., and Pandis, S. N.: High formation of secondary organic
  aerosol from the photo-oxidation of toluene, Atmospheric Chemistry and Physics, 9,
  2973-2986, 2009.
- Hoffmann, T., Odum, J. R., Bowman, F., Collins, D., Klockow, D., Flagan, R. C., and
  Seinfeld, J. H.: Formation of organic aerosols from the oxidation of biogenic
  hydrocarbons, Journal of Atmospheric Chemistry, 26, 189-222,
  10.1023/a:1005734301837, 1997.
- 960 Kanakidou, M., Seinfeld, J. H., Pandis, S. N., Barnes, I., Dentener, F. J., Facchini, M. C.,

961 Van Dingenen, R., Ervens, B., Nenes, A., Nielsen, C. J., Swietlicki, E., Putaud, J. P.,

962 Balkanski, Y., Fuzzi, S., Horth, J., Moortgat, G. K., Winterhalter, R., Myhre, C. E. L.,

963 Tsigaridis, K., Vignati, E., Stephanou, E. G., and Wilson, J.: Organic aerosol and global

964 climate modelling: a review, Atmospheric Chemistry and Physics, 5, 1053-1123, 2005.

- Karnezi, E., Riipinen, I., and Pandis, S. N.: Measuring the atmospheric organic aerosol
  volatility distribution: a theoretical analysis, Atmospheric Measurement Techniques, 7,
  2953—2965, 2014.
- Keywood, M. D., Varutbangkul, V., Bahreini, R., Flagan, R. C., and Seinfeld, J. H.:
  Secondary organic aerosol formation from the ozonolysis of cycloalkenes and related
  compounds, Environmental Science & Technology, 38, 4157-4164, 10.1021/es0353630,
  2004.
- 972 Kokkola, H., Yli-Pirila, P., Vesterinen, M., Korhonen, H., Keskinen, H., Romakkaniemi,

973 S., Hao, L., Kortelainen, A., Joutsensaari, J., Worsnop, D. R., Virtanen, A., and Lehtinen,

974 K. E. J.: The role of low volatile organics on secondary organic aerosol formation,

- 975 Atmospheric Chemistry and Physics, 14, 1689-1700, 10.5194/acp-14-1689-2014, 2014.
- Kostenidou, E., Pathak, R. K., and Pandis, S. N.: An algorithm for the calculation of
  secondary organic aerosol density combining AMS and SMPS data, Aerosol Sci.
  Technol., 41, 1002-1010, 10.1080/02786820701666270, 2007.
- Wall Partitioning in Teflon Environmental Chambers Using Rapid Bursts of Low-

- Volatility Oxidized Species Generated in Situ, Environmental Science & Technology, 50,
  5757-5765, 10.1021/acs.est.6b00606, 2016.
- Kristensen, K., Cui, T., Zhang, H., Gold, A., Glasius, M., and Surratt, J. D.: Dimers in
  alpha-pinene secondary organic aerosol: effect of hydroxyl radical, ozone, relative
  humidity and aerosol acidity, Atmospheric Chemistry and Physics, 14, 4201-4218,
  10.5194/acp-14-4201-2014, 2014.
- Kroll, J. H., Chan, A. W. H., Ng, N. L., Flagan, R. C., and Seinfeld, J. H.: Reactions of
  semivolatile organics and their effects on secondary organic aerosol formation,
  Environmental Science & Technology, 41, 3545-3550, 10.1021/es062059x, 2007.
- 990 Kuwata, M., and Martin, S. T.: Phase of atmospheric secondary organic material affects 991 Acad. Sci. U. S. 109. 17354-17359, its reactivity. Proc. Natl. A., 992 10.1073/pnas.1209071109, 2012.
- La, Y. S., Camredon, M., Ziemann, P. J., Valorso, R., Matsunaga, A., Lannuque, V., LeeTaylor, J., Hodzic, A., Madronich, S., and Aumont, B.: Impact of chamber wall loss of
  gaseous organic compounds on secondary organic aerosol formation: explicit modeling
  of SOA formation from alkane and alkene oxidation, Atmos. Chem. Phys., 16, 14171431, 10.5194/acp-16-1417-2016, 2016.
- Loza, C. L., Chan, A. W. H., Galloway, M. M., Keutsch, F. N., Flagan, R. C., and
  Seinfeld, J. H.: Characterization of Vapor Wall Loss in Laboratory Chambers,
  Environmental Science & Technology, 44, 5074-5078, 10.1021/es100727v, 2010.
- Loza, C. L., Chhabra, P. S., Yee, L. D., Craven, J. S., Flagan, R. C., and Seinfeld, J. H.:
  Chemical aging of m-xylene secondary organic aerosol: laboratory chamber study,
  Atmospheric Chemistry and Physics, 12, 151-167, 10.5194/acp-12-151-2012, 2012.
- Matsunaga, A., and Ziemann, P. J.: Gas-Wall Partitioning of Organic Compounds in a
  Teflon Film Chamber and Potential Effects on Reaction Product and Aerosol Yield
  Measurements, Aerosol Sci. Technol., 44, 881-892, 10.1080/02786826.2010.501044,
  2010.

- 1008 McMurry, P. H., and Grosjean, D.: Gas and Aerosol Wall Losses in Teflon Film Smog
- 1009 Chambers, Environmental Science & Technology, 19, 1176-1182, 10.1021/es00142a006,
  1010 1985.
- 1011 McMurry, P. H., and Rader, D. J.: Aerosol Wall Losses in Electrically Charged 1012 Chambers, Aerosol Sci. Technol., 4, 249-268, 10.1080/02786828508959054, 1985.
- 1013 McVay, R. C., Cappa, C. D., and Seinfeld, J. H.: Vapor-Wall Deposition in Chambers:
- 1014 Theoretical Considerations, Environmental Science & Technology, 48, 10251-10258,
  1015 10.1021/es502170j, 2014.
- 1016 McVay, R. C., Zhang, X., Aumont, B., Valorso, R., Camredon, M., La, Y. S., Wennberg,
- 1017 P. O., and Seinfeld, J. H.: SOA formation from the photooxidation of  $\alpha$ -pinene:
- 1018 systematic exploration of the simulation of chamber data, Atmos. Chem. Phys., 16, 2785-
- 1019 2802, 10.5194/acp-16-2785-2016, 2016.
- Miles, R. E. H., Reid, J. P., and Riipinen, I.: Comparison of Approaches for Measuring
  the Mass Accommodation Coefficient for the Condensation of Water and Sensitivities to
  Uncertainties in Thermophysical Properties, J. Phys. Chem. A, 116, 10810-10825,
  10.1021/jp3083858, 2012.
- 1024 Ng, N. L., Kroll, J. H., Keywood, M. D., Bahreini, R., Varutbangkul, V., Flagan, R. C., 1025 Seinfeld, J. H., Lee, A., and Goldstein, A. H.: Contribution of first- versus second-1026 generation products to secondary organic aerosols formed in the oxidation of biogenic 1027 Technology, 40, hydrocarbons, Environmental Science & 2283-2297, 1028 10.1021/es052269u, 2006.
- 1029 Ng, N. L., Kroll, J. H., Chan, A. W. H., Chhabra, P. S., Flagan, R. C., and Seinfeld, J. H.:
- 1030 Secondary organic aerosol formation from m-xylene, toluene, and benzene, Atmospheric
- 1031 Chemistry and Physics, 7, 3909-3922, 2007.
- 1032 Odum, J. R., Hoffmann, T., Bowman, F., Collins, D., Flagan, R. C., and Seinfeld, J. H.:
- 1033 Gas/Particle Partitioning and Secondary Organic Aerosol Yields, Environmental Science
- 1034 & Technology, 30, 2580-2585, 10.1021/es950943+, 1996.
- Odum, J. R., Jungkamp, T. P. W., Griffin, R. J., Flagan, R. C., and Seinfeld, J. H.: The
  atmospheric aerosol-forming potential of whole gasoline vapor, Science, 276, 96-99,
  10.1126/science.276.5309.96, 1997a.
- Odum, J. R., Jungkamp, T. P. W., Griffin, R. J., Forstner, H. J. L., Flagan, R. C., and
  Seinfeld, J. H.: Aromatics, reformulated gasoline, and atmospheric organic aerosol
  formation, Environmental Science & Technology, 31, 1890-1897, 10.1021/es9605351,
  1041 1997b.
- Pathak, R. K., Presto, A. A., Lane, T. E., Stanier, C. O., Donahue, N. M., and Pandis, S.
  N.: Ozonolysis of alpha-pinene: parameterization of secondary organic aerosol mass
  fraction, Atmospheric Chemistry and Physics, 7, 3811-3821, 2007a.
- 1045 Pathak, R. K., Stanier, C. O., Donahue, N. M., and Pandis, S. N.: Ozonolysis of alpha-
- 1046 pinene at atmospherically relevant concentrations: Temperature dependence of aerosol
- 1047 mass fractions (yields), J. Geophys. Res.-Atmos., 112, 8, 10.1029/2006jd007436, 2007b.
- 1048 Perraud, V., Bruns, E. A., Ezell, M. J., Johnson, S. N., Yu, Y., Alexander, M. L.,
- 1049 Zelenyuk, A., Imre, D., Chang, W. L., Dabdub, D., Pankow, J. F., and Finlayson-Pitts, B.
- 1050 J.: Nonequilibrium atmospheric secondary organic aerosol formation and growth, Proc.
- 1051 Natl. Acad. Sci. U. S. A., 109, 2836-2841, 10.1073/pnas.1119909109, 2012.
- Pierce, J. R., Engelhart, G. J., Hildebrandt, L., Weitkamp, E. A., Pathak, R. K., Donahue,
  N. M., Robinson, A. L., Adams, P. J., and Pandis, S. N.: Constraining particle evolution
  from wall losses, coagulation, and condensation-evaporation in smog-chamber
  experiments: Optimal estimation based on size distribution measurements, Aerosol Sci.
  Technol., 42, 1001-1015, 10.1080/02786820802389251, 2008.
- Presto, A. A., Hartz, K. E. H., and Donahue, N. M.: Secondary organic aerosol
  production from terpene ozonolysis. 2. Effect of NOx concentration, Environmental
  Science & Technology, 39, 7046-7054, 10.1021/es050400s, 2005.
- Presto, A. A., and Donahue, N. M.: Investigation of alpha-pinene plus ozone secondary
  organic aerosol formation at low total aerosol mass, Environmental Science &
  Technology, 40, 3536-3543, 10.1021/es052203z, 2006.

- Pye, H. O. T., Chan, A. W. H., Barkley, M. P., and Seinfeld, J. H.: Global modeling of
  organic aerosol: the importance of reactive nitrogen (NOx and NO3), Atmospheric
  Chemistry and Physics, 10, 11261-11276, 10.5194/acp-10-11261-2010, 2010.
- 1066 Renbaum-Wolff, L., Grayson, J. W., Bateman, A. P., Kuwata, M., Sellier, M., Murray, B.

1067 J., Shilling, J. E., Martin, S. T., and Bertram, A. K.: Viscosity of alpha-pinene secondary

1068 organic material and implications for particle growth and reactivity, Proc. Natl. Acad.

- 1069 Sci. U. S. A., 110, 8014-8019, 10.1073/pnas.1219548110, 2013.
- Saha, P. K., and Grieshop, A. P.: Exploring Divergent Volatility Properties from Yield
  and Thermodenuder Measurements of Secondary Organic Aerosol from α-Pinene
  Ozonolysis, Environmental Science & Technology, 50, 5740-5749,
  10.1021/acs.est.6b00303, 2016.
- Saleh, R., Donahue, N. M., and Robinson, A. L.: Time Scales for Gas-Particle
  Partitioning Equilibration of Secondary Organic Aerosol Formed from Alpha-Pinene
  Ozonolysis, Environmental Science & Technology, 47, 5588-5594, 10.1021/es400078d,
  2013.
- Saukko, E., Lambe, A. T., Massoli, P., Koop, T., Wright, J. P., Croasdale, D. R.,
  Pedernera, D. A., Onasch, T. B., Laaksonen, A., Davidovits, P., Worsnop, D. R., and
  Virtanen, A.: Humidity-dependent phase state of SOA particles from biogenic and
  anthropogenic precursors, Atmospheric Chemistry and Physics, 12, 7517-7529,
  10.5194/acp-12-7517-2012, 2012.
- Saunders, S. M., Jenkin, M. E., Derwent, R. G., and Pilling, M. J.: Protocol for the
  development of the Master Chemical Mechanism, MCM v3 (Part A): tropospheric
  degradation of non-aromatic volatile organic compounds, Atmos. Chem. Phys., 3, 161180, 10.5194/acp-3-161-2003, 2003.
- Seinfeld, J. H., and Pandis, S. N.: Atmospheric chemistry and physics : from air pollution
  to climate change, 2nd ed., Wiley, Hoboken, N.J., xxviii, 1203 p. pp., 2006.
- Shilling, J. E., Chen, Q., King, S. M., Rosenoern, T., Kroll, J. H., Worsnop, D. R.,
  McKinney, K. A., and Martin, S. T.: Particle mass yield in secondary organic aerosol

- 1091 formed by the dark ozonolysis of alpha-pinene, Atmospheric Chemistry and Physics, 8,1092 2073-2088, 2008.
- 1093 Shilling, J. E., Chen, Q., King, S. M., Rosenoern, T., Kroll, J. H., Worsnop, D. R.,
- 1095 dependent elemental composition of alpha-pinene SOA particles, Atmospheric Chemistry

DeCarlo, P. F., Aiken, A. C., Sueper, D., Jimenez, J. L., and Martin, S. T.: Loading-

1096 and Physics, 9, 771-782, 2009.

- 1097 Shiraiwa, M., and Seinfeld, J. H.: Equilibration timescale of atmospheric secondary 1098 organic aerosol partitioning, Geophys. Res. Lett., 39, 6, 10.1029/2012gl054008, 2012.
- 1099 Song, C., Zaveri, R. A., Alexander, M. L., Thornton, J. A., Madronich, S., Ortega, J. V.,
- 1100 Zelenyuk, A., Yu, X. Y., Laskin, A., and Maughan, D. A.: Effect of hydrophobic primary
- 1101 organic aerosols on secondary organic aerosol formation from ozonolysis of alpha-
- 1102 pinene, Geophys. Res. Lett., 34, 5, 10.1029/2007gl030720, 2007.
- Stanier, C. O., Pathak, R. K., and Pandis, S. N.: Measurements of the volatility of
  aerosols from alpha-piniene ozonolysis, Environmental Science & Technology, 41, 27562763, 10.1021/es0519280, 2007.
- 1106 Tabge, O. GNU Parallel The Command-Line Power Tool, :login: The USENIX
  1107 Magazine, <u>http://www.gnu.org/s/parallel</u>, 2011.
- 1108 Tsigaridis, K., Daskalakis, N., Kanakidou, M., Adams, P. J., Artaxo, P., Bahadur, R.,
- 1109 Balkanski, Y., Bauer, S. E., Bellouin, N., Benedetti, A., Bergman, T., Berntsen, T. K.,
- 1110 Beukes, J. P., Bian, H., Carslaw, K. S., Chin, M., Curci, G., Diehl, T., Easter, R. C.,
- 1111 Ghan, S. J., Gong, S. L., Hodzic, A., Hoyle, C. R., Iversen, T., Jathar, S., Jimenez, J. L.,
- 1112 Kaiser, J. W., Kirkevåg, A., Koch, D., Kokkola, H., Lee, Y. H., Lin, G., Liu, X., Luo, G.,
- 1113 Ma, X., Mann, G. W., Mihalopoulos, N., Morcrette, J. J., Müller, J. F., Myhre, G.,
- 1114 Myriokefalitakis, S., Ng, N. L., O'Donnell, D., Penner, J. E., Pozzoli, L., Pringle, K. J.,
- 1115 Russell, L. M., Schulz, M., Sciare, J., Seland, Ø., Shindell, D. T., Sillman, S., Skeie, R.
- 1116 B., Spracklen, D., Stavrakou, T., Steenrod, S. D., Takemura, T., Tiitta, P., Tilmes, S.,
- 1117 Tost, H., van Noije, T., van Zyl, P. G., von Salzen, K., Yu, F., Wang, Z., Wang, Z.,
- 1118 Zaveri, R. A., Zhang, H., Zhang, K., Zhang, Q., and Zhang, X.: The AeroCom evaluation

- and intercomparison of organic aerosol in global models, Atmos. Chem. Phys., 14,
  10845-10895, 10.5194/acp-14-10845-2014, 2014.
- Vaden, T. D., Song, C., Zaveri, R. A., Imre, D., and Zelenyuk, A.: Morphology of mixed
  primary and secondary organic particles and the adsorption of spectator organic gases
  during aerosol formation, Proc. Natl. Acad. Sci. U. S. A., 107, 6658-6663,
  10.1073/pnas.0911206107, 2010.
- Vaden, T. D., Imre, D., Beranek, J., Shrivastava, M., and Zelenyuk, A.: Evaporation
  kinetics and phase of laboratory and ambient secondary organic aerosol, Proc. Natl.
  Acad. Sci. U. S. A., 108, 2190-2195, 10.1073/pnas.1013391108, 2011.
- 1128 Virtanen, A., Joutsensaari, J., Koop, T., Kannosto, J., Yli-Pirila, P., Leskinen, J., Makela,
- 1129 J. M., Holopainen, J. K., Poschl, U., Kulmala, M., Worsnop, D. R., and Laaksonen, A.:
- 1130 An amorphous solid state of biogenic secondary organic aerosol particles, Nature, 467,
- 1131 824-827, 10.1038/nature09455, 2010.
- Virtanen, A., Kannosto, J., Kuuluvainen, H., Arffman, A., Joutsensaari, J., Saukko, E.,
  Hao, L., Yli-Pirila, P., Tiitta, P., Holopainen, J. K., Keskinen, J., Worsnop, D. R., Smith,
  J. N., and Laaksonen, A.: Bounce behavior of freshly nucleated biogenic secondary
  organic aerosol particles, Atmospheric Chemistry and Physics, 11, 8759-8766,
  10.5194/acp-11-8759-2011, 2011.
- Weitkamp, E. A., Sage, A. M., Pierce, J. R., Donahue, N. M., and Robinson, A. L.:
  Organic aerosol formation from photochemical oxidation of diesel exhaust in a smog
  chamber, Environmental Science & Technology, 41, 6969-6975, 10.1021/es070193r,
  2007.
- Yeh, G. K., and Ziemann, P. J.: Alkyl Nitrate Formation from the Reactions of C-8-C-14
  n-Alkanes with OH Radicals in the Presence of NOx: Measured Yields with Essential
  Corrections for Gas-Wall Partitioning, J. Phys. Chem. A, 118, 8147-8157,
  10.1021/jp500631v, 2014.
- Yeh, G. K., and Ziemann, P. J.: Gas-Wall Partitioning of Oxygenated OrganicCompounds: Measurements, Structure-Activity Relationships, and Correlation with Gas

- 1147 Chromatographic Retention Factor, Aerosol Sci. Technol., 49, 726-737,
  1148 10.1080/02786826.2015.1068427, 2015.
- 1149 Ye, P., Ding, X., Hakala, J., Hofbauer, V., Robinson, E. S., and Donahue, N. M.: Vapor
- 1150 wall loss of semi-volatile organic compounds in a Teflon chamber, Aerosol Sci. Technol.,
- 1151 50, 822-834, 10.1080/02786826.2016.1195905, 2016.
- Zaveri, R. A., Easter, R. C., Shilling, J. E., and Seinfeld, J. H.: Modeling kinetic
  partitioning of secondary organic aerosol and size distribution dynamics: representing
  effects of volatility, phase state, and particle-phase reaction, Atmospheric Chemistry and
  Physics, 14, 5153-5181, 10.5194/acp-14-5153-2014, 2014.
- Zhang, X., Pandis, S. N., and Seinfeld, J. H.: Diffusion-Limited Versus QuasiEquilibrium Aerosol Growth, Aerosol Sci. Technol., 46, 874-885,
  10.1080/02786826.2012.679344, 2012.
- Zhang, X., Cappa, C. D., Jathar, S. H., McVay, R. C., Ensberg, J. J., Kleeman, M. J., and
  Seinfeld, J. H.: Influence of vapor wall loss in laboratory chambers on yields of
  secondary organic aerosol, Proc. Natl. Acad. Sci. U. S. A., 111, 5802-5807,
  10.1073/pnas.1404727111, 2014.
- 1163 Zhang, X., Schwantes, R. H., McVay, R. C., Lignell, H., Coggon, M. M., Flagan, R. C.,
- and Seinfeld, J. H.: Vapor wall deposition in Teflon chambers, Atmospheric Chemistry and Physics, 15, 4197-4214, 10.5194/acp-15-4197-2015, 2015a.
- 1166 Zhang, X., McVay, R. C., Huang, D. D., Dalleska, N. F., Aumont, B., Flagan, R. C., and 1167 Seinfeld, J. H.: Formation and evolution of molecular products in alpha-pinene secondary А., 1168 organic aerosol. Proc. Natl. Acad. Sci. U. S. 112, 14168-14173, 1169 10.1073/pnas.1517742112, 2015b.
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1173	<b>Table 1:</b> Experimental conditions and results for the $\alpha$ -pinene ozonolysis experiments					
	Experiment	Initial Seed	Initial	$\Delta M_o^{b}$	SOA Mass	
	-	Surface Area	[α-pinene] <sup>a</sup>	$(\mu g m^{-3})$	Yield <sup>c</sup> (%)	
		$(\mu m^2 cm^{-3})$	$(\mu g m^{-3})$			
	100 ppb O <sub>3</sub> nucleation	0	290.2±23.2	$62.0\pm1.2^{d}$	22.6±1.9	
	100 ppb $O_3$ low AS	1130	280.5±22.4	$63.0 \pm 0.8^{d}$	23.3±1.9	
	100 ppb O <sub>3</sub> high AS	2700	238.7±19.1	$50.6 \pm 1.6^{d}$	23.3±1.9	
	500 ppb $O_3$ nucleation	0	274.4±21.9	$87.3 \pm 0.3^{e}$	31.8±2.5	
	500 ppb O <sub>3</sub> low AS	1300	264.9±21.2	$75.7 \pm 0.6^{e}$	28.6±2.3	
	500 ppb O <sub>3</sub> high AS	2720	236.1±18.9	$66.3 \pm 1.9^{e}$	28.1±2.4	
1174	<sup>a</sup> Concentration of $\alpha$ -pinene injected into the chamber. All the $\alpha$ -pinene reacted in the 500					
1175	ppb $O_3$ experiments, but not the 100 ppb $O_3$ experiments.					
1176	<sup>b</sup> Uncertainties in the peak SOA mass concentration ( $\Delta M_o$ ) are calculated from one					
1177	standard deviation of the aerosol volume as measured by the scanning mobility particle					
1178	sizer.					
1179	<sup>c</sup> SOA mass yields at peak SOA growth are reported.					
1180	<sup>d</sup> The SOA mass concentration is calculated using the density = $1.39 \text{ g cm}^{-3}$ obtained from					
1181	the 100 ppb $O_3$ nucleation experiment.					
1182	<sup>e</sup> The SOA mass concentration is calculated using the density = $1.37 \text{ g cm}^{-3}$ obtained from					
1183	the 500 ppb $O_3$ nucleation e			.)		
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**Table 1:** Experimental conditions and results for the  $\alpha$ -pinene ozonolysis experiments

<b>Table 2.</b> Coupled vapor-particle dynamics model parameters						
Definition	Value					
Vapor-particle mass accommodation coefficient	0.1					
Vapor-wall mass accommodation coefficient	10-6					
Timescale of oligomerization	4 h					
Branching ratios and saturation concentrations of	$[0.6 (>10^3), 0.3(10^2),$					
oxidation products	0.05(10), 0.05(1) and					
	0(0.1)]					
Gas-phase molecular diffusivity	$\frac{0(0.1)]}{3 \times 10^{-6} \text{ m}^2 \text{ s}^{-1}}$ 2.5 m <sup>-1</sup>					
Surface area-to-volume ratio of the chamber						
Equivalent organic mass concentration in the wall	$10 \text{ mg m}^{-3}$					
Eddy diffusion coefficient	$0.03 \text{ s}^{-1}$					
Molecular weight of the diffusing gas-phase molecule i	168, 184, 192, 200 and					
	$216 \text{ g mole}^{-1}$					
Initially absorbing organic material in seed aerosol	0.01 µg m <sup>-3</sup>					
Pressure	$1 \times 10^5$ Pa					
Temperature	298 K					
Density of inorganic seed	1700 kg m <sup>-3</sup>					
Density of organic material on seed particle	1300 kg m <sup>-3</sup>					
	DefinitionVapor-particle mass accommodation coefficientVapor-wall mass accommodation coefficientTimescale of oligomerizationBranching ratios and saturation concentrations of oxidation productsGas-phase molecular diffusivitySurface area-to-volume ratio of the chamberEquivalent organic mass concentration in the wallEddy diffusion coefficientMolecular weight of the diffusing gas-phase molecule iInitially absorbing organic material in seed aerosolPressureTemperatureDensity of inorganic seed					

**Table 2:** Coupled vapor-particle dynamics model parameters



**Figure 1:** Particle wall deposition coefficients ( $\beta_i$ ) measured during the low AS-seed only and high AS-seed only experiments in GTEC. Also shown are the particle wall deposition coefficients (labeled "Other") measured in previous routine monthly AS-seed only experiments in the chamber. These previous routine monthly AS-seed only experiments were performed using either a 0.008 M AS or a 0.1 M AS solution. Coagulation-corrected particle wall deposition coefficients (see Pierce et al. (2008) and main text for details) are also shown, using dashed lines.



Figure 2: Reaction profiles of the  $\alpha$ -pinene ozonolysis experiments. Panels (a), (b) and 1214 1215 (c) show results from the nucleation, low AS and high AS 100 ppb O<sub>3</sub> experiments, 1216 respectively. Panels (d), (e) and (f) show results from the nucleation, low AS and high AS 1217 500 ppb O<sub>3</sub> experiments, respectively. As explained in the main text, the SOA mass 1218 concentrations ( $\Delta M_0$ ) for the nucleation and low AS experiments are obtained using the 1219 particle wall deposition rates obtained from the low AS-seed only experiments, while the 1220 SOA mass concentrations ( $\Delta M_0$ ) for the high AS-seed experiments are obtained using the 1221 particle wall deposition rates obtained from the high AS-seed only experiments.



**Figure 3:** Time-dependent SOA growth curves for α-pinene ozonolysis. Panels (a) and 1225 (b) show 10 min-averaged results from the 100 ppb and 500 ppb  $O_3$  experiments, 1226 respectively. Only SOA growth data up to the point of SOA peak growth are shown.



1229 **Figure 4:** 10 min-averaged SOA mass yields over the course of an  $\alpha$ -pinene ozonolysis 1230 experiment as a function of initial total AS seed surface area concentration for the (a) 100 1231 ppb O<sub>3</sub> experiments, and (b) 500 ppb O<sub>3</sub> experiments. Symbol color indicates the SOA mass concentration and symbol size indicates the time after O<sub>3</sub> is injected into the 1232 chamber. The imes symbols are the SOA mass yields at peak SOA growth obtained from 1233 1234 the experimental data. The y-axis error bars represent the uncertainty in the SOA mass 1235 yield at peak SOA growth, which originates from the  $\alpha$ -pinene injection and the aerosol 1236 volume concentration measured by the SMPS at peak SOA growth (one standard 1237 deviation).



1241 Figure 5: Comparison of SOA mass yields obtained in this study to those of previous 1242 dark  $\alpha$ -pinene ozonolysis studies (Table S3). The SOA mass yields and concentrations of 1243 majority of these previous studies (Hoffmann et al., 1997; Griffin et al., 1999; Cocker et 1244 al., 2001b; Gao et al., 2004; Presto et al., 2005; Presto et al. 2006; Pathak et al., 2007b; 1245 Song et al., 2007) were previously compiled by Shilling et al. (2008). Similar to Shilling 1246 et al. (2008), all the data shown here (including those reported in this study) have been adjusted using an organic density of 1.0 g cm<sup>-3</sup>, and to 298 K using a temperature 1247 correction of 1.6 % per K, as recommended by Pathak et al. (2007b) to facilitate easier 1248 1249 comparison among the different studies.



- **Figure 6:** SOA mass yields at peak SOA growth as a function of both the seed surface area and O<sub>3</sub> concentration for  $\alpha_p = 1$ , 0.1, 0.01, and 0.001. The SOA mass yields at peak SOA growth are indicated by colors and contours. Note that the color bars for panels (a), (b) and (c) have different SOA mass yield ranges. Simulations were carried out using the optimal branching ratios, oligomerization rate, and vapor wall deposition rate parameters obtained in this study. The initial  $\alpha$ -pinene concentration was set to 50 ppb, and a fixed O<sub>3</sub> concentration was used in place of a linear injection.
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