1 For a point by point answer to each question of the reviewers, please see our

2 published comment at ACPD. It contains a respond to each question and

3 suggestion the two reviewers raised. Changes to the manuscript are as follows:

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Evaluating Secondary Inorganic Aerosols in 3-Dimensions

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13 Abstract.

14 The spatial distribution of aerosols and their chemical composition dictates whether aerosols 15 have a cooling or a warming effect on the climate system. Hence, properly modeling the 3-16 dimensional distribution of aerosols is a crucial step for coherent climate simulations. Since 17 surface measurement networks only give 2-D data, and most satellites supply integrated 18 column information, it is thus important to integrate aircraft measurements in climate model 19 evaluations. In this study, the vertical distribution of secondary inorganic aerosol (i.e. sulfate, 20 ammonium and nitrate) is evaluated against a collection of 14 AMS flight campaigns and surface measurements from 2000-2010 in the USA and Europe. GISS ModelE2 is used with 21 22 multiple aerosol microphysics (MATRIX, OMA) and thermodynamic (ISORROPIA II, EQSAM) 23 configurations. Our results show that the MATRIX microphysical scheme improves the model 24 performance for sulfate, but that there is a systematic underestimation of ammonium and 25 nitrate over the USA and Europe in all model configurations. In terms of gaseous precursors, 26 nitric acid concentrations are largely underestimated at the surface while overestimated in the 27 higher levels of the model, Heterogeneous reactions on dust surfaces is an important sink for 28 nitric acid, even high in the troposphere. At high altitudes, nitrate formation is calculated to be

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ammonia limited. The underestimation of ammonium and nitrate in polluted regions is most likely caused by a too simplified treatment of the NH_3/NH_4^+ partitioning which affects the HNO_3/NO_3^- partitioning.

34 1. Introduction

35 The impact of aerosols on climate and air quality is a function of their chemical composition, 36 abundance and spatial distribution. Understanding the vertical profile of aerosols is crucial for 37 radiative forcing calculations (Xu and Penner, 2012), since aerosols interact with radiation 38 directly through absorption and scattering (Bauer and Menon, 2012; Haywood and Boucher, 39 2000; Stocker et al., 2013), and indirectly via interactions with clouds (Lohmann and Feichter, 40 2005). Comparisons of model results with organic aerosol aircraft data showed large 41 discrepancies in the free troposphere (Heald et al., 2005, 2011). Sulfate and ammonium nitrate 42 aerosols, although much simpler to model than organics, have not been studied in the vertical 43 in much detail. There is large uncertainty in the magnitude of the forcing induced by sulfate and 44 ammonium nitrate aerosols, with estimates for the preindustrial to present day direct radiative forcing of sulfate ranging from -0.6 to -0.2 Wm⁻² while for ammonium nitrate from -0.3 to -0.03 45 Wm⁻² (Stocker et al., 2013) under present day conditions. These forcings are projected to 46 47 change in the future, driven by trends in precursor emissions. The projected increase in 48 agricultural ammonia emissions, which will result in greater availability of ammonia, contrasted 49 with the projected reductions in NO_x emissions, can lead to an increased relative contribution 50 of ammonium nitrate to the total secondary inorganic aerosol (SIA) abundance, due to the 51 strong projected decrease of sulfate aerosols (Hauglustaine et al., 2014; Hodas et al., 2014). 52 Yet, the effect of these changes on ammonium nitrate concentrations are still a matter of active

research: *Paulot et al.* [2016] showed increases in nitrate load in the free troposphere, while surface concentrations decreased, and *Pusede et al.* [2016] showed changes in tropospheric chemistry in western USA, with increased ammonium nitrate production during daytime rather than at night.

Thermodynamically, ammonia tends to neutralize sulfuric acid over the highly volatile nitric acid (Tagaris et al., 2007). The formation of fine-mode nitrate is a function of ammonia, sulfate availability and relative humidity (RH), since its precursor, nitric acid, condenses following thermodynamic equilibrium (Potukuchi and Wexler, 1995a, 1995b). Sulfuric acid and nitric acid also participate in heterogeneous uptake on dust particles, forming coarse sulfate and nitrate, a process that acts as a sink for the gas phase precursors (Bauer and Koch, 2005; Ravishankara, 1997).

64 In this paper we evaluate ammonium, nitrate and sulfate aerosols in the NASA GISS ModelE2 65 against surface and aircraft observations, extending what Bauer et al. [2007b] did for nitrate 66 aerosol for the year 2000, by using new aerosol configurations that had been implemented in 67 GISS ModelE2 since then, and a substantially extended record of SIA measurements, both from 68 ground stations and various flight campaigns. To assess the model in terms of SIA surface 69 distribution and vertical profiles, we evaluated the performance of three aerosol 70 configurations, described in section 2.1.1, by comparing them against surface data measured 71 over the USA and Europe during 2000-2010, and 14 flight campaigns, as described in section 72 2.2. We then study the climatology of the model against measurements, both at surface and at 73 higher altitudes (sections 3.1-3.3), and explore the model uncertainties with the help of 74 sensitivity experiments (section 3.4).

75 2. Experimental approach

76 2.1 Model description

77 The NASA GISS ModelE2 model (Schmidt et al., 2014) was run with interactive tropospheric 78 (Shindell et al., 2001, 2003) and stratospheric chemistry (Shindell et al., 2006) and coupled with 79 three different aerosol configurations, as described below. A horizontal resolution of 2° in 80 latitude by 2.5° in longitude and a vertical resolution of 40 layers to 0.1 hPa was used. The 81 simulation was nudged using 6 hourly National Centers for Environmental Prediction (NCEP) 82 reanalysis data (Kalnay et al., 1996) for the horizontal wind component. Sea surface 83 temperatures (SSTs) and sea ice cover were prescribed using the Met Office Hadley Center's sea ice and sea surface temperature data set (HadISST1) (Rayner et al., 2003). 84

The nitrate optical depth of GISS ModelE2 in the CMIP5 archive was found to be problematic, consistent with the findings of *Shindell et al.* [2013] for a likely too high nitrate load. <u>The model</u> was using the Henry value of ammonia instead of the effective Henry value, which resulted in large abundances of ammonia, hence ammonium, hence nitrate. In our work the nitrate scheme had been corrected and nitrate distribution in the column reflects surface sources such as agricultural, industrial and biomass burning areas.

91 2.1.1 Aerosols schemes

Two aerosol schemes were used in this study: OMA (One Moment Aerosol) (Koch et al., 2006; Miller et al., 2006) and MATRIX (Multiconfiguration Aerosol TRacker of mlXing state) (Bauer et al., 2008). OMA is a bulk mass scheme with one fine mode bin of prescribed size for SO_4^{2-} , NH_4^+ , and NO_3^- . In OMA, heterogeneous uptake of SO_2 and HNO_3 on dust surfaces is also included, which takes place on the three smallest size bins out of the five size bins used for mineral dust

97 (Bauer et al., 2004, 2007). This was changed after Bauer et al., 2007 where dust was 98 represented in four size classes, and coating on all classes was tracked. MATRIX is a 99 microphysical scheme representing nucleation, condensation and coagulation. Sulfate is 100 tracked with both number and mass concentrations for 16 populations, which are based on 101 mixing state. MATRIX represents an intermediate level of complexity; only the total mass of 102 nitrate, ammonium and aerosol water is calculated, and then distributed across populations 103 based on the sulfate abundance in each one of them, assuming internally mixed components. 104 This approach greatly reduces the required number of transported variables.

Due to the focus on SIA in this paper we will give a brief description of the sulfate and nitrate schemes in our model. The sulfate chemistry module in both schemes, OMA and MATRIX, is based on Koch et al. (1999) and includes prognostic calculation of gas and aqueous phase DMS, MSA, SO₂ and sulfate concentrations. This provides the sulfate mass in the OMA scheme, and provides aqueous sulfate production rates and H₂SO₄ concentrations as input parameters for MATRIX microphysics (Bauer et al., 2008).

111 To partition between the gas and particle phases the model uses the non-linear 112 thermodynamics. Both schemes were run coupled to the secondary inorganic aerosol 113 thermodynamics scheme EQSAM (Metzger et al., 2002a, 2002b). MATRIX was also run coupled 114 to ISORROPIA II (Fountoukis and Nenes, 2007), which was only recently introduced into GISS 115 ModelE2. EQSAM is a parameterized thermodynamics scheme that relies on the relationship 116 between activity coefficients and RH to calculate the solute activity and the non-ideal solution 117 properties, while ISORROPIA II calculates the equilibrium constants and solves the 118 thermodynamic equations analytically. Both models use the same input parameters: NH_x

119 $(NH_3+NH_4^+)$, $SO_4^{2^-}$, XNO_3 $(HNO_3+NO_3^-)$, RH and temperature, and interactively calculate the $SO_4^{2^-}$ 120 , NH_4^+ , NO_3^- and aerosol H_2O concentrations at equilibrium, as well as the residual NH_3 and 121 HNO_3 in the gas phase.

The thermodynamical equilibrium for Aitken mode sized particles, which is important for CCN, might not be properly captures by models (*Benduhn et al.* [2016]). This is not expected to be a problem in this study because Aitken mode particles are a small fraction of the total aerosol mass. In addition, for the coarse mode, large uncertainties exist regarding the availability of crustal and coarse mode material in equilibrium thermodynamic calculations. Our simulations do not take into consideration crustal (e.g. Mg²⁺, K⁺, Ca²⁺) and sea salt (e.g. Na⁺, Cl⁻) ions in the thermodynamics, although this option is available in the model.

129 The model ran in the following three configurations: OMA-EQSAM, MATRIX-EQSAM, and 130 MATRIX-ISORROPIA, and we are comparing model $PM_{2.5}$ (particles with dry diameter smaller 131 than 2.5 μ m) with measured $PM_{2.5}$ at surface, and model PM_1 (particles with dry diameter 132 smaller than 1 μ m) with measured PM_1 at the vertical, for consistency with the available 133 measurements.

134 **2.1.2 Emissions**

This study used the Coupled Model Intercomparison Project phase 5 (CMIP5) historical anthropogenic emissions until 2005 (Lamarque et al., 2010) and the Representative Concentration Pathway 4.5 (RCP4.5) scenario thereafter (van Vuuren et al., 2011). Biomass burning emissions came from the Global Fire Emissions Database (GFED3) inventory (van der Werf et al., 2010). The emissions include seasonal variations for the biomass burning, soil NO_x, shipping and aircraft sectors (Lamarque et al., 2010), yet lack seasonal variability for all other

anthropogenic emissions, including agricultural NH₃ sources. In order to prevent unrealistic ammonium and nitrate aerosols loads during wintertime, the agricultural NH₃ emissions were altered using the local solar zenith angle, in order to produce a more realistic seasonal variability, but kept the total annual emissions the same. This approach is comparable to *Adams et al.* [1999] and *Park* [2004] who scaled ammonia emissions from crops and fertilizers according to the number of daylight hours.

147 **2.1.3 Sensitivity runs**

148 NH₃ emissions are controlled by the agricultural sector (Lamarque et al., 2010), both in the USA 149 and Europe, where more than 80% of NH₃ emissions are agriculture-related (van Damme et al., 150 2015; Paulot et al., 2014). We test how changing agricultural NH₃ emissions affect ammonium 151 nitrate formation under two scenarios: double and five times higher agricultural NH₃ emissions, 152 using the MATRIX-ISORROPIA aerosol configuration. The results of that sensitivity are presented 153 in section 3.4.

154 2.2 Observational datasets

155 **2.2.1 Surface measurements**

We evaluate our simulations against nitrate and sulfate PM_{2.5} data measured by the Interagency Monitoring of Protected Visual Environments (IMPROVE) network over the continental United States (Malm et al., 1994, 2004), and against ammonia, ammonium, nitric acid, nitrate, SO₂ and sulfate measured by the European Monitoring and Evaluation Programme (EMEP), available via the NILU-EBAS database, for the years 2000-2010. From EMEP we use the corrected sulfate for sea salt (XSO₄) (EMEP, 2014, Chp. 3) as it better represents fine sulfate.

162 IMPROVE currently has 212 sites, predominantly rural (Hand et al., 2011, 2012), while EMEP has 163 around 40 sites measuring aerosol composition in Europe, many of which are urban (Tørseth et al., 2012). The data in Europe is reported in μ gX m⁻³ (where X is either sulfur or nitrogen) and in 164 the USA in μ g m⁻³. We decided to keep these units unchanged in the rest of the manuscript, and 165 166 convert the units of the model to represent those of the measurements, rather than doing the 167 opposite. We compared monthly mean values from all available stations with monthly mean 168 model output. An examination of the mean spatial distribution over the USA (Figure 1) revealed 169 distinct regimes with different pollution levels, which motivated a regional division of the data 170 into eastern USA (EUSA) and western USA (WUSA). Europe (EU; Figure 2) and the Arctic (ARC; 171 data from flight campaigns only) were studied independently (Table 1). The standard deviation, 172 correlation coefficient (R), and normalized mean bias (NMB) between the monthly mean 173 surface values within the studied regions (black frames in Figures 1 and 2) and the model's 174 monthly mean at the stations locations in each region, were calculated. It is important to note 175 that during the 11-year period the number of measuring sites has varied in each region, and 176 that not all stations measured all species.

177 2.2.2 Flight campaigns

The Aerodyne Aerosol Mass Spectrometer (AMS), which measures chemical composition and size distribution of non-refractory particles (such as ammonium, nitrate and sulfate) with diameter smaller than 1 μ m (Allan et al., 2003; Jimenez et al., 2003), had been part of many flight campaigns in the past decade. Another common method to measure inorganic particle composition is using the particle-into-liquid-sampler (PILS), which quantifies the ionic content of particulate matter using ion chromatography (Weber et al., 2001). In this study we use data

184 from 14 flight campaigns, two of which used the PILS instrument for chemical composition 185 measurements, and the rest used the AMS (Table 2). The flights took place in the Northern 186 Hemisphere during short campaign periods, predominantly during spring and summer seasons, 187 between 2001-2011. The flight tracks of the campaigns used here are presented in Figure 3. 188 Data were retrieved using the Tools for Airborne Data interface (https://tad.larc.nasa.gov/), as 189 well as the AMS global database (https://sites.google.com/site/amsglobaldatabase/). For every 190 campaign a mean regional vertical profile was calculated by averaging the flight data within the 191 model's grid. For short-range campaigns such as ACE, CRISTAL, MILAGRO, TexAQS, and EUCAARI 192 all available data were used, for ITOP the transit flight data were parsed out, and for the rest of 193 the campaigns only data within the regional boundaries we study (black frames in Figure 3) 194 were used. These boundaries were chosen in accordance with the surface observations.

195 The campaign-average profile was compared against the monthly mean model output, a not 196 uncommon practice in model-aircraft comparison studies (e.g. Bauer et al., 2007; Emmons et 197 al., 2000; Shindell et al., 2003). The simulations were sub-sampled by taking into consideration 198 the geographical variability of the flights, but not the sub-monthly temporal variability, to yield 199 a mean corresponding profile. The one standard deviation variability of the campaign data per 200 model level was calculated for the measurements and model simulations, which represents the 201 spatial variability of the concentrations during the whole field campaign for the measurements, 202 and the spatial variability of the monthly mean modeled concentrations for the model. The 203 duration of the field campaigns ranged from 7 to 17 days. In the Results section we picked 4 204 representative campaigns that display systematic behavior, one for each region (Figure 7). The 205 rest of the campaigns can be found in the appendix (Figure A2, A3).

206 3. Results and discussion

In terms of mean surface concentrations (measured and modeled) in the Western Hemisphere sulfate concentrations are higher than nitrate concentrations. That is not the case in the Eastern Hemisphere, since over western Europe sulfate and nitrate aerosols are comparable in mass (Figure 2), consistent with *Schaap et al.* [2004]. At the whole atmospheric column (not shown here), sulfate peaks over east EU and northern Africa due to in-cloud production and transport, while the nitrate column distribution corresponds to the surface distribution, with maxima over the continental hot spots, driven by urban pollution and biomass burning.

214 **3.1 Surface climatology**

215 Surface data show high concentrations of nitrate and sulfate in the industrialized EUSA and EU 216 and lower concentrations in WUSA, with some urban hot spots (Figures 1 and 2). We compared 217 the model skill, with respect to measurements, under the three different aerosol configurations 218 in Figure 4 for nitrate (left) and sulfate (right). The regional clusters observed reflect the fact 219 that performance in terms of R and NMB is controlled by region rather than aerosol scheme. 220 For sulfate, the simulation with no microphysics (OMA, blue) is always biased lower (by 1-4%) 221 compared to the other two simulations (MATRIX, red and green). This result is due to the 222 microphysical processes included in MATRIX (i.e. nucleation, condensation and coagulation), 223 which allow for aerosols to spread over the entire size distribution, including the existence of 224 smaller particles (the freshly nucleated ones), which sediment more slowly. Additionally the 225 solubility of sulfate in MATRIX is calculated as a weighted average of the mixed particle 226 component's solubility (SO4 mixed with dust, BC, etc.), and is always less than the pure SO4 227 solubility in OMA. The differences in both size and solubility lead to a longer lifetime of MATRIX

Keren Mezuman 5/25/2016 12:56 PM Deleted:), Keren Mezuman 5/25/2016 12:56 PM Deleted: a result Keren Mezuman 5/25/2016 12:53 PM Deleted: of having sulfate aerosols spread over the entire size distribution,

232 sulfate, thus increasing the aerosol mass. As an indication, the mean lifetime of sulfate in 2005 233 was 4.2 days in the two MATRIX simulations, against 3.2 days in the OMA simulation. We observe a systematic underestimation of ammonium, nitrate and sulfate in EUSA and EU (35% 234 235 for nitrate, 30% for ammonium, 20% for sulfate). Despite the negative bias, the three aerosol 236 types correlate well with measurements in these regions (R>0.5). This high correlation is due to 237 the fact the simulations successfully capture the aerosol seasonal cycle (discussed in the next 238 section). In the WUSA, the simulations overestimate sulfate by 12%, and underestimated 239 nitrate by 80%, while there is no correlation between the model and observations for nitrate. 240 The different behavior across regions reflects the fact that the WUSA is driven by agricultural 241 emissions while in the EUSA industrial and residential emissions dominate. The ability of the 242 model to capture the seasonality (discussed in the next section) is important for model skill and 243 is discussed in the next section.

244 **3.2 Surface seasonality**

245 Figure 5 shows that in the EU there is little variation in the SO₂ seasonality between the three 246 simulations, which is emission-level driven. The modeled surface concentration overestimates measurements by about 0.5 μ gS m⁻³, with an amplified seasonal cycle. Past studies (Dentener et 247 248 al., 2006; Vestreng et al., 2007) have raised concerns regarding the accuracy of SO₂ emission 249 inventories, which might be part of the explanation of the SO₂ overestimation. Additionally, 250 wintertime chemistry slowdown due to reduced photochemistry increases the SO₂ lifetime, 251 resulting in reduced sulfate formation rates, contributing to the underestimation of sulfate 252 concentration which can be as high as a factor of 2 during winter months. For sulfate, the 253 difference between the simulations is dominated by the aerosol scheme, with the summertime

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Keren Mezuman 5/25/2016 12:57 PM Deleted: leading to longer lifetime of sulfate mass in MATRIX. 256 peak being more pronounced in the MATRIX simulations than in the OMA one. As explained in 257 the previous section, MATRIX simulates higher concentrations due to the existence of smaller particles with longer lifetimes compared to OMA. Surface NH₃ (Figure 5) is overestimated in all 258 259 three simulations, which might be due to incorrect NH_x partitioning calculated by EQSAM and 260 ISORROPIA II, a hypothesis that is supported by the underestimate of ammonium. Contrary to 261 SO₂ and NH₃, nitric acid is underestimated by the simulations by a factor of 3. This contributes 262 to the underestimation of nitrate in all simulations. The simulated seasonality of nitrate 263 matches that of the measurements, peaking during winter and reaching a minimum during 264 summer. Konovalov et al. [2008] identified a slight underestimation of NO_x in emission 265 inventories in southern Europe, which would contribute to underestimations of XNO₃.

266 IMPROVE has extensive sulfate and nitrate surface data to compare against the model 267 simulations. EMEP provides additional HNO₃ data from 9 stations, predominantly around the 268 Great Lakes, which is not enough for a proper regional analysis. Unfortunately, ammonium and 269 gas phase aerosol precursors are not routinely measured via the IMPROVE network. In the 270 eastern USA (Figure 6) the model simulations exhibit peak sulfate concentrations during 271 summer, with the MATRIX simulations having a stronger seasonality than OMA, which better 272 matches observations. For nitrate, all simulations systematically underestimate measurements 273 during most of the year (by about 0.2 μ g m⁻³), except during winter, where MATRIX slightly 274 overestimates them (less then $0.1 \,\mu g \, m^{-3}$). The HNO₃ underestimation by the model, as evident 275 by the limited measurements we obtained in EUSA (Figure A1), contributes to the nitrate 276 underestimation.

277 In WUSA the simulated sulfate and nitrate seasonality (Figure 6 left panels) is flat compared to the measurements. For sulfate, the measured range is 0.7 μ g m⁻³, while in the MATRIX 278 simulations the range is 0.25 μ g m⁻³ and OMA-EQSAM is 0.15 μ g m⁻³. All simulations 279 280 underestimate measurements during summer and overestimate them during winter. The 281 measured maximum sulfate concentrations are around summer. This feature is captured by 282 OMA-EQSAM, but the MATRIX simulations calculate spring and fall peaks instead. For nitrate, 283 the measurements peak in early winter, a feature that is not captured by the simulations, as 284 modeled nitrate peaks in winter. During the winter OMA-EQSAM and MATRIX EQSAM are 285 similar, probably due to the common thermodynamical scheme, while MATRIX-ISORROPIA II is 286 higher by 0.05 μ g m⁻³. Modeled nitrate is underestimated compared to measurements 287 throughout the year: in the MATRIX simulations it is underestimated by about 0.45 μ g m⁻³ (80%) 288 of the measured value), and in OMA-ISORROPIA it is underestimated by about 0.4 μ g m⁻³.

289 3.3 Vertical Profiles

290 The simulated mean vertical profiles of sulfate, ammonium, nitric acid (when available), and 291 nitrate are evaluated against the mean measured profiles in Figure 7. The measured and 292 modeled standard deviations (gray shading and dashed lines, respectively), along with the 293 number of days each layer was sampled (black squares), are shown as well. Generally, aerosol 294 concentrations decrease with altitude as they peak near emission sources at the surface. Some 295 of the data used in this study were affected by intense fire plumes (Fisher et al., 2010; Jacob et 296 al., 2010), as can be seen in the ATCPAC (ARC) and ARCTAS spring and summer (ARC) panels 297 (Figure 7 and A2). Fires act as a source of NO_x , NH_3 and SO_2 , increasing the concentration of

sulfate, ammonium and nitrate in the measurements. Fire emissions are included in our simulations, yet these emissions could be underestimated, as *lchoku and Ellison* [2014] indicated is the case in many bottom-up emission inventories such as GFED3 (used here), and are also a function of properly resolving the transport. Even if all these factors are accurate in the model, the monthly mean output we use would dilute the signal of a fire event as observed in a flight profile.

304 Modeled sulfate concentrations are underestimated compared to the measurements (first 305 column in Figure 7 and Figures A2 and A3). The MATRIX simulations that include aerosol 306 microphysics show higher concentrations compared to the bulk scheme. During INTEX-A (EUSA) 307 the MATRIX simulations produced in the boundary layer around 1 μ g m⁻³ higher sulfate 308 concentrations compared to OMA. The thermodynamic scheme (EQSAM or ISORROPIA II) 309 makes a minor difference for sulfate, stemming from the simulations' climate feedbacks, with 310 the green and red lines overlaying each other. All these results are consistent with the ones 311 presented earlier for the surface.

In remote environments like the Florida Keys (CRISTAL-FACE, Figure A3), Azores (ITOP-UK, Figure A3) and the Arctic (ARCTAS spring and summer, Figure 7 and A2), ammonium and nitrate concentrations are generally very low, and the models are able to reproduce the aerosol concentrations. However, in campaigns over land such as: EUCAARI EU, EUSA: INTEX-A, NEAQS, DISCOVER-MD, CALNEX WUSA, TexAQS, and Mexico: MILAGRO-MIRAGE, INTEX-B, there is consistent underestimation of both ammonium and nitrate, especially in the boundary layer

318 (Figures 7, A2 and A3). The sensitivity runs we performed, presented later, explore whether this

319 is due to precursor levels or to the thermodynamic parameterization used.

320 From the nitric acid profiles (third column in Figure 7 and Figure A2), it is evident that the 321 model strongly overestimates the measurements in the middle and upper troposphere. HNO_3 322 overestimation at high altitudes is consistent with Figure 10 from Shindell et al. [2006], On top 323 of that, the modeled nitric acid shows distinct OMA and MATRIX profiles, which diverge with 324 increasing height, with differences that can become as high as 0.3 ppbv. Though there is not 325 much dust at these altitudes, the inclusion of heterogeneous reactions on dust surfaces in OMA 326 is the main difference in the gas phase chemistry of OMA and MATRIX schemes. The coarse 327 mode nitrate mass formed by those heterogeneous reactions almost fully accounts for the 328 difference in HNO₃ between the two schemes. However, this loss is insufficient to explain the 329 discrepancy between the model and measurements. We exclude the nitrate that forms on dust 330 (coarse nitrate) from the nitrate profiles, since neither they are in the PM1 aerosol 331 measurements, nor they are calculated in the MATRIX simulations.

The overestimation of nitric acid does not result in overestimation of nitrate, which is also affected by the availability of both sulfate and ammonia, on top of environmental factors like relative humidity and temperature. Even though nitrate concentrations are low in many locations (below 2 μ g m⁻³), the simulations underestimates it to be below 0.1 μ g m⁻³ in EUSA (INTEX-A in Figure 7, NEAQS, DISCOVER-MD, and TexAQS in Figure A2), WUSA (CALNEX in Figure 7), arctic (ARCPAC in Figure A2), Central America (INTEX-B in Figure A2, MILAGRO-MIRAGE, and CRISTAL-FACE in Figure A3), consistent with the spring-summer surface underestimation.

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Deleted: , who attributed this model overestimation to enhanced stratosphere- troposphere exchange, which brings HNO ₃ - rich air marcor from the lower stratosphere to

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the upper troposphere

347 Another key point is that there is little difference in the nitrate concentrations simulated by the 348 different aerosol configurations. Differences between the simulations are evident only in the boundary layer in EUCAARI (EU, Figure 7), ~0.8 μ g m⁻³, and ACE-ASIA (Japan, Figure A3), ~0.3 μ g 349 350 m⁻³. In these locations, the difference is not evident on a thermodynamic scheme basis, but 351 rather on a microphysical scheme, with MATRIX-EQSAM and MATRIX-ISORROPIA grouped 352 against OMA-EQSAM. The difference in concentration between the simulations is also evident in the ammonium profiles of these campaigns. In EUCAARI, nitrate and ammonium have higher 353 354 concentrations in the OMA-EQSAM simulation, while sulfate is consistently larger in the 355 MATRIX ones. In ACE-ASIA however, both sulfate and ammonium concentrations are higher 356 with OMA-EQSAM, yet nitrate concentrations are higher in the MATRIX simulations. It is 357 evident from these profiles that the simulations with lower sulfate concentrations are also the 358 simulations with higher nitrate concentrations. The role of thermodynamics to the NH_3/NH_4^+ 359 partitioning at different NH₃ levels will be discussed in the next section.

360 3.4 Sensitivity runs

In order to study the interplay between precursor concentrations and thermodynamics we perturbed the ammonia emissions from agriculture. For these runs, presented in Figure 8, we use the MATRIX-ISORROPIA scheme with standard NH₃ emission (green line), double agricultural NH₃ emissions (purple line) and five times agricultural NH₃ emissions (brown line). At the surface, as NH₃ emissions are increased, the ammonium and nitrate underestimation by the model disappears (Figure 8). However, a comparison with the limited available surface NH₃ measurements reveals that even with the standard NH₃ emissions the model overestimates

368 NH₃ concentrations. This is also evident in TexAQS and CALNEX (WUSA) NH₃ profiles (Figure A4). 369 Similarly, in the vertical, with increasing NH₃ emissions the nitric acid model overestimation 370 decreases (Figure A5), as more NH₃ becomes available to react with nitric acid and partition it 371 to the aerosol phase. These results indicate that the NH_3/NH_4^+ partitioning is not accurately 372 calculated by the model, and that this strongly affects the nitric acid/nitrate partitioning. 373 Further evidence to support our conclusion lies in Figure 9 that presents the modeled and 374 measured partitioning ratios (NH₃ over total NH_x, and HNO₃ over total XNO₃). For NH_x all three 375 simulations are grouped together, while for XNO₃ a distinct difference between the 376 thermodynamic schemes is revealed: MATRIX-EQSAM overestimate the partitioning ratio 377 during the summer, and MATRIX-ISORROPIA II is closer to measurements. From the surface 378 seasonality of the individual species (Figure 5) it is clear that the divergence in the ratio is 379 driven mainly by nitrate concentrations, as HNO₃ concentrations are the same for MATRIX-380 EQSAM and MATRIX-ISORROPIA II (red and green curves overlaying each other). The difference 381 between these two simulations in terms of nitrate concentrations are of the order of 0.05 μ g N 382 m^{-3} and are most distinct during summer (Figure 5). Similarly, the difference between the 383 simulations for XNO3 is greater during summer. Thermodynamically, other than precursor 384 levels, the difference in behavior in summer and the rest of the year is also controlled by 385 temperature and RH.

386 Conclusions

In this work we used a collection of surface measurements and flight campaigns over the USA
 and Europe from 2000-2010 to study the regional and vertical distribution of secondary

inorganic aerosols and their precursors under different aerosol configurations of the GISS
 ModelE2. In the USA sulfate aerosol dominate the near surface SIA composition, but over EU
 the nitrate aerosol contribution is comparable in mass.

392 We compare the behavior of SIA concentrations in high (EUSA, EU) and low (WUSA) aerosol 393 precursor source regions, as the relative contribution of different sectors generates different 394 chemical regimes. We observe a systematic underestimation of near surface concentrations in 395 the EUSA and EU compared to the surface network measurements: 35% for nitrate, 30% for 396 ammonium, 20% for sulfate. However, despite the negative bias, all three simulations have high 397 correlation coefficients (R>0.5) when compared against surface data. In the WUSA the results 398 for sulfate and nitrate are different in sign, sulfate is biased high (12%) with R=0.43, while 399 nitrate is biased low (80%) with no correlation between the simulations and the measurements 400 (R<0.1). The correlation is also driven by the difficulty of the model to capture the annual cycle 401 of the species.

402 Microphysics has improved the sulfate simulation, as the MATRIX scheme yields consistently, 403 both at the surface and in the vertical, higher sulfate concentrations, due to smaller particles 404 having longer lifetimes compared to OMA, the bulk scheme (4.2 days against 3.2 days). For 405 ammonium nitrate simulations there is an additional level of complexity in the form of accurate 406 thermodynamics, which is sensitive both to the precursors and to environmental parameters such as temperature and humidity. Since we have performed nudged simulations, they do not 407 408 show big differences in temperature and RH, so the differences between the simulations are 409 expected to be dominated by the thermodynamical scheme and not the underlying

meteorological parameters. In terms of precursors, NH₃ is slightly overestimated, as indicated
by surface measurements over EU in Figure 5 and TexAQS and CALNEX campaigns in Figure A3.
HNO₃ is underestimated at the surface but overestimated at higher levels, and jncluding
heterogeneous reactions on dust surfaces decreases the overestimation. A more complex
version of MATRIX, that currently does not exist, should include heterogeneous uptake on dust.

415 Overall, aerosol mass is consistently underestimated both at surface and in the boundary layer.

416 In our sensitivity runs, increasing NH₃ emissions results in NH₃ overestimation, however it 417 improves our simulated HNO₃ profiles. When more NH₃ is available it reacts with HNO₃ to form 418 ammonium nitrate, resolving underestimations in the aerosol phase. Hence, the partitioning of 419 NH_x which strongly affects the partitioning of XNO₃ is not accurately simulated in the model. 420 Aan de Brugh et al. [2012] identified an overestimation of gas phase precursors during daytime 421 (equivalent to summer) and overestimation of aerosol phase species during nighttime 422 (equivalent to winter), and found it to be related to the time scale of vertical mixing against the 423 timescale of thermodynamic equilibrium. This relationship was not analyzed here, since it 424 requires high temporal resolution model output.

An examination of aerosol pH (not presented here) indicated a pH rage from 1 to 2 over EU. This range was recently identified by *Weber et al.* [2016] as a buffering pH zone where partitioning of ammonium nitrate between the gas and aerosol phases is sensitive. Thus, ions which affect pH might play an important role in nitrate formation. Hence, taking into consideration crustal and sea salt ions could affect our thermodynamics and partitioning in regions where these ions are abundant, as *Karydis et al.* [2016] demonstrated. However, these Keren Mezuman 6/20/2016 10:43 PM **Deleted:** .

Keren Mezuman 6/20/2016 10:43 PM **Deleted:** The overestimation is attributed to too strong stratosphere-troposphere exchange, yet

are currently tracked as bulk dust and sea salt aerosols in the model. In addition to tracking Na⁺,

436 Cl⁻, etc. separately, we would need to consider the different time scales of the thermodynamics

437 associated with aerosol size distribution. In the future, we plan to investigate the influence of

438 pH on the results in more detail.

439 In this paper we have demonstrated the importance of size resolved sulfate chemistry. 440 However, currently we treat nitrate as bulk, as it is computationally expensive to add 15 nitrate 441 tracers. Perhaps underestimation of nitrate is not only a matter of thermodynamics but 442 microphysics as well, and that properly resolving the size distribution, and considering the 443 chemistry that depends on that would improve our simulations. 444 Acknowledgements. Climate modeling at GISS is supported by the NASA Modeling, Analysis, 445 and Prediction program. Resources supporting this work were provided by the NASA High-End 446 Computing (HEC) Program through the NASA Center for Climate Simulation (NCCS) at Goddard

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Table 1. Regional boundaries for data division

Region	Boundaries
Arctic (ARC)	55°-90°N, 60°-170°W
Eastern USA (EUSA)	30°-50°N, 60°-95°W
Western USA (WUSA)	30°-50°N, 114°-130°W
European Union (EU)	35°-70°N, 10°W-30°E

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Table 2. Airborne measurements used in this stud	iy.
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Campaign (Aircraft)	Region (season, year)	Technique and reference	Regime
ACE-Asia (CIPRAS TWIN OTTER)	Japan (spring, 2001)	AMS (Huebert, 2003)	Polluted
CRYSTAL- FACE (CIPRAS TWIN OTTER)	South Florida (summer, 2002)	AMS (Conant et al., 2004)	Polluted
ITOP (BAE- 146)	Azores (summer. 2004)	AMS (Fehsenfeld et al., 2006)	Remote
INTEX-A (DC- 8, J-31)	Eastern USA (summer, 2004)	CIMS (HNO ₃), PILS (SO ₄ ,NH ₄ ,NO ₃) (Singh et al., 2006)	Polluted
NEAQS (NOAA-P3)	Eastern USA (summer, 2004)	CIMS (HNO ₃), AMS (SO ₄ , NH ₄ , NO ₃) (Fehsenfeld et al., 2006)	Polluted
INTEX-B (DC- 8)	Western USA (spring, 2006)	CIMS (HNO ₃), AMS (SO ₄ ,NH ₄ ,NO ₃) (Leaitch et al., 2009)	Polluted
MILAGRO (C120)	Mexico (spring, 2006)	AMS (DeCarlo et al., 2008)	Polluted
TexAQS (NOAA-P3)	Texas (fall, 2006)	CIMS (NH3, HNO ₃), AMS (SO ₄ , NH ₄ , NO ₃) (Parrish et al., 2009)	Polluted

EUCAARI (BAE-146)	NW EU (spring, 2008)	AMS (Morgan et al., 2010)	Polluted
ARCPAC (NOAA-P3	Arctic (spring, 2008)	CIMS (HNO ₃), AMS (SO ₄ , NH ₄ , NO ₃) (Fisher et al., 2010)	Fire
ARCTAS (DC- 8, P-3)	Arctic (spring/summer 2008)	CIMS (HNO ₃), AMS (SO ₄ , NH ₄ , NO ₃) (Jacob et al., 2010)	Fire
CALNEX (NOAA P-3)	West coast (summer, 2010)	CIMS (HNO ₃ , NH ₃), AMS (SO ₄ , NH ₄ , NO ₃) (Ryerson et al., 2013)	Polluted
DISCOVER- MD (P-3B, UC-12)	Maryland (summer, 2011)	TD-LIF (HNO_3) (Anderson et al., 2014), PILS (SO ₄ , NH_4 , NO_3) (Ziemba et al., 2013)	Polluted



779 2010) simulated by MATRIX-EQSAM overlaid by measurements from the IMPROVE network. 780

781 The model data units match the units of the measured data ($\mu g m^{-3}$).



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sulfate).



- **Figure 3.** Flight tracks of 14 flight campaigns used in this study (2001-2011).





Figure 4. Surface regional statistics (2000-2010). Left panel: nitrate and ammonium (data

- available only for EU); right panel: sulfate. The correlation coefficient (R) between the
- rgg simulation and the measurements is in the y-axis, and normalized mean bias (NMB) is in the x-
- 800 axis. MATRIX-EQSAM is in red, MATRIX-ISORROPIA II is in green and OMA-EQSAM is in blue.



Surface EU (2000-2010)



804 Figure 5. 2000-2010 mean annual cycle over Europe, error bars represent ± one standard

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deviation. Measurements are in orange, MATRIX-EQSAM is in red, MATRIX-ISORROPIA II is in
 green and OMA-EQSAM is in blue.





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808 **Figure 6.** 2000-2010 mean annual cycle over WUSA (left) and EUSA (right), error bars represent

± one standard deviation. Measurements are in orange, MATRIX-EQSAM is in red, MATRIX 810 ISORROPIA II is in green and OMA-EQSAM is in blue.



Figure 7. Mean regional concentration profiles from the arctic (first row), eastern USA (second

row), western USA (third row) and Europe (fourth row). First column is SO₄, second is NH₄, third
 is HNO₃ and fourth is NO₃.



817 Figure 8. 2000-2010 mean annual cycle over Europe, error bars represent ± one standard

818 deviation. Measurements are in orange, MATRIX-ISORROPIA II: with regular emissions is in

 $green, with \ double \ agricultural \ NH_3 \ emissions \ is \ in \ purple, \ and \ with \ 5-times \ agricultural \ NH_3$

emissions is in brown.

Surface ratio EU (2000-2010)





822 **Figure 9.** 2000-2010 mean partitioning ratio annul cycle over Europe, error bars represent ±

- 823 one standard deviation. Measurements are in orange, MATRIX-EQSAM is in red, MATRIX-
- 824 ISORROPIA II is in green and OMA-EQSAM is in blue.







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Figure A2. Mean regional concentration profiles. First column is SO₄, second is NH₄, third is



 $$\rm HNO_3$ and fourth is NO_3. Measurements are in orange, MATRIX-EQSAM is in red, MATRIX-

⁸³⁴ ISORROPIA II is in green and OMA-EQSAM is in blue.



837 Figure A2: continued



839 Figure A3. Mean regional concentration profiles. First column is SO₄, second is NH₄, and third is

840 NO₃. Measurements are in orange, MATRIX-EQSAM is in red, MATRIX-ISORROPIA II is in green

⁸⁴¹ and OMA-EQSAM is in blue.







845 is in green, with double agricultural NH₃ emissions is in purple, and with 5-times agricultural

846 NH_3 emissions is in brown.







- $849 \qquad \text{Figure A5.} \ \text{Mean regional HNO}_3 \ \text{profiles from the arctic, EUSA and WUSA.} \ \text{Measurements are in}$
- 850 orange, MATRIX-ISORROPIA II: with regular emissions is in green, with double agricultural NH₃
- 851 emissions is in purple, and with 5-times agricultural NH₃ emissions is in brown.