July 26, 2017

Dear Editor,

We would like to thank you for serving as the editor for our manuscript. We have revised the manuscript to incorporate all of the general and specific comments and suggestions from the two reviewers. As the reviewers have suggested, we have removed section 3.5 and some sentences in other sections that appeared speculative, where we only had weak evidence to justify our arguments. Similarly, we rephrased several sentences and made them more understandable by adding further information. We believe that with these revisions the interpretation of the findings and hence the quality of the manuscript has improved significantly.

Sincerely yours,

Khadak Singh Mahata on behalf of all coauthors

Seasonal and diurnal variations of methane and carbon dioxide in the Kathmandu Valley in the foothills of the central Himalaya

by Khadak Singh Mahata et al., 2017 (ACPD)

We would like to thank both reviewers for their constructive comments and suggestions. Please find below the reviewers' comments in black, our response in blue and changes in the revised manuscript in red. The line numbers in our response refer to the line numbers in the revised manuscript.

Reviewer 1

General comments: Atmospheric greenhouse gases (GHGs) such as CO2, CH4, H2O and CO are important climate forcing agents having significant impacts on climate system and air quality. This study brings outs first continuous measurements of atmospheric GHGs using high precision cavity ring down spectrometer (Picarro 24G2401, USA) at Kathmandu Valley during March 2013 to March 2014. The authors have done and extensive study on GHGs variability with time and space. However, there are a few minor technical changes in the manuscript. This paper is recommend to publish in ACP after incorporating the minor technical corrections.

We would like to thank reviewer for considering that the manuscript contains extensive work on the variability of GHGs in the Kathmandu Valley. We have incorporated the reviewer's comments and suggestions to the extent possible in the following sections of the manuscript.

Line18-20: This paper studies about GHGs and GHGs are not classified as pollutants especially CO2 and CH4. Impact of pollution on GHGs need to be emphasis not to refer GHGs as pollutants. Also sentence "This paper reports …". May be re-written

We agree with the reviewer that CO_2 and CH_4 at normal ambient concentrations are not considered as pollutants from the health perspective. Thus, we did not describe them as pollutants in the manuscript. We slightly modified the sentence "this paper reports…" to

distinguish the GHGs from the pollutant CO, and thus more understandably reflect what has been looked into in this study.

Line51-53: Not clear. Authors may please check the sentence. "All three species showed strong diurnal and saying immediately CH4 and CO did not show any variation…". May be provided quantitative numbers.

Thank you for pointing out that the statement was not clear to the reviewer. Here we are explaining the diurnal variations of CH_4, CO_2 and CO at two measurement sites (Bode and Chanban). We have articulated it by rephrasing the sentence as follows (lines 52-56):

At Bode, all three gas species $(CO_2, CH_4$ and CO) showed strong diurnal patterns in their mixing ratios with a pronounced morning peak (ca. 08:00), a dip in the afternoon, and again gradual increase through the night until the next morning, whereas CH_4 and CO mixing ratios at Chanban did not show any noticeable diurnal variations.

Line139: Rupakheti et al., 2016 need to be updated if available

The Rupakheti et al. manuscript will be submitted to ACPD soon. We have updated the citation as Rupakheti et al. (2017, manuscript in preparation) in line 144-145.

Line252: Sentence "The % may be written as Difference (%) of the analyzer differed by…".

Corrected. The sentence has been rephrased as follows (lines 258-260):

The difference between CO_2 mixing ratio reported by the analyzer and the reference mixing ratio was within 5%.

Line349: Units of GHGs and other gases should be uniform in the manuscript.

Corrected. They are now reported in ppm throughout the manuscript.

Line395-396: Impact of rainfall on CO2 dilution process may be supported with reference Mahesh et al., 2014 "Impact of land-sea breeze…".

Frequent rainfall suppresses emission sources which results in reducing mixing ratios of the gas and aerosol species. To reflect this, we have modified the sentences in lines 414-422:

The concentrations of most pollutants in the region are lower during the monsoon period (Sharma et al., 2012, Marinoni, 2013; Putero et al., 2015) because frequent and heavy rainfall suppresses emission sources. We saw a drop in the CO_2 mixing ratio during the rainfall period due to changes in various processes such as enhanced vertical mixing, uptake of CO_2 by vegetation and soils, and where relevant reduction in combustion sources. CO_2 can also dissolve into rainfall, forming carbonic acid, which may lead to a small decrease in the CO_2 mixing ratio as has been observed during heavy intensity rainfall (Mahesh et al., 2014; Chaudhari et al., 2007).

The following references are added in the list of references.

Mahesh P, Sharma N, Dadhwal VK, Rao PVN, Apparao BV, et al.: Impact of Land-Sea Breeze and Rainfall on CO2Variations at a Coastal Station. J. Earth Sci. Clim. Change, 5:201. doi: 10.4172/2157-7617.1000201, 2014.

Chaudhari, P.R., Gajghate, D.G., Dhadse, S. et al. Environ. Monit. Assess., 135: 281. doi:10.1007/s10661-007-9649-7, 2007.

Line432-433: Please check the statement that CO2 will be high but CH4 will be high during post-monsoon season.

This comment does not match with the lines mentioned by the reviewer. We did not find the exact or similar sentence in the manuscript.

Figure 4: Titled should be changed. Since GHGs are not pollutants and legend should be CH4 not Ch4.

In order to remove confusion, we have changed the caption of Figure 4 as follows:

Relation between mixing ratios and wind direction observed at Bode in the Kathmandu Valley (a) CH_4, (b) CO_2, and (c) CO from March 2013 to February 2014. The figure shows variations of CH_4, CO_2 and CO mixing ratios based on frequency counts of wind direction (in %) as represented by circle. The color represents the different mixing ratios of the gaseous species. The units of CH_4, CO_2 and CO are in ppm.

The legend is also corrected as suggested.

Figure11: Show double Y-axis for better visualization

Thank you for the suggestion. As individual species varies over a wide range and there is also an order of magnitude difference in their mixing ratios, putting them in double Y-axes with different scales makes the figure more confusing. We have therefore kept the figure as it is.

Reviewer 2

Highly precise and long-term measurements of greenhouse gases are essential to understand the underlying processes in the context of global climate change. It is particularly valuable in regions like Asia where it is currently limited by dedicated long term observations. This study demonstrates the observational variations of CO2 and CH4 mixing ratios in the Kathmandu Valley (Nepal), and compares them to those of other rural sites. The scope of this study is hence highly relevant to the public and authors' efforts on this regard need to be appreciated. However, the manuscript in its present form does not meet the standard to merit the publication in ACP, and needs to be adequately revised. I therefore recommend the current manuscript to undergo major revision to be considered in ACP.

We would like to thank reviewer for considering our measurements as important for the region and that the scope of our study is relevant to the public. We have revised the manuscript as suggested by the reviewer.

General comments

As mentioned above, I value authors' effort in this study as an important step towards generating observational wealth which can be tremendously used by scientific community to understand a wide range of mechanisms involved in this aspect. Given the complicated interplay of many processes involved, single approach cannot answer the unresolved scientific questions related to these processes and mechanisms affecting mixing ratio variations. This requires defining far more systematic approaches and other robust tools. However the finding from this study can be very valuable if presented with adequate measurement and analysis methods/techniques used, including a good summary of the methods and a much clear report on the observational variations of these analysed tracers. The results will be more convincing by focusing on the key aspects of the data rather than trying to relate them to flux categories based on assumption (many times) and without sufficient tools (inverse modeling). In some places, results are presented nicely, but then an explanation is suggested based on other literature that focussed on other study regions/methods, without even demonstrating its relevance to this dataset. Also in some places, the study gets into overly ambitious interpretations of the results and conclusions on the basis of single analysis. On this basis, I recommend authors to focus on presenting this study by clearly stating the measurement and analysis techniques used, defining their strategies, providing analysed results & possible uncertainties and giving more convincing interpretations of the results and conclusions. I highly recommend authors not to jump to interpreting emission/flux sources and patterns based on the single site measurements and the analysis done in this study. Sections in this manuscript dealing with these aspects needs to be (preferably) removed or highly restructured.

Thank you for the general comments. We would like to break down the general comments, as suggested by the editor, to address reviewer's general comments systematically and clearly as follows:

As mentioned above, I value authors' effort in this study as an important step towards generating observational wealth which can be tremendously used by scientific community to understand a wide range of mechanisms involved in this aspect.

Thanks for considering our work as observational wealth that can be useful for the scientific community to understand the process involved in this aspect.

Given the complicated interplay of many processes involved, single approach cannot answer the unresolved scientific questions related to these processes and mechanisms affecting mixing ratio variations. This requires defining far more systematic approaches and other robust tools. However the finding from this study can be very valuable if presented with adequate measurement and analysis methods/techniques used, including a good summary of the methods and a much clear report on the observational variations of these analysed tracers. The results will be more convincing by focusing on the key aspects of the data rather than trying to relate them to flux categories based on assumption (many times) and without sufficient tools (inverse modeling).

We agree with this comment that a single approach in our study cannot answer the resolved scientific questions. We apologize for not mentioning the approaches clearly. We have improved measurement and analysis methods as: (i) resolved the confusion between uncertainty and variability (expressed as standard deviation) and recalculated annual values of all species from monthly data to make them comparable with background and regional sites, as explained in the specific comments below (also refer to lines 316-322 in the manuscript), (ii) performed significance test of CH4 and CO2 data between Bode and Mauna Loa, and Bode and Waliguan. Similarly, in the results part, we have removed (i) the loose connection to regional pollution in section 3.5 (influence of regional emission and transport) and (ii) substantially revised section 3.6 (CO and CO2 ratio: potential emission sources), clearly mentioning that the information on sources inferred from the ratio analysis is only indicative rather than definitive (lines 680-682 and lines 715-730) and (iii) also included useful statistics in Table 6.

In some places, results are presented nicely, but then an explanation is suggested based on other literature that focussed on other study regions/methods, without even demonstrating its relevance to this dataset.

We have selected other urban/semi-urban sites in India as mentioned in the manuscript on the basis of similarity of emission sources and similar influence of regional meteorology in the region (as mentioned in lines 325-328). The background sites (Mauna Loa and Waliguan) were selected to see the spatial difference between the background sites and current study sites as described in lines 322-325 in the manuscript.

Also in some places, the study gets into overly ambitious interpretations of the results and conclusions on the basis of single analysis.

We would like to thank reviewer for noticing these ambitious interpretations. We have rephrased such interpretations, mainly in the lines 715-730 and lines 835-837.

On this basis, I recommend authors to focus on presenting this study by clearly stating the measurement and analysis techniques used, defining their strategies, providing analysed results & possible uncertainties and giving more convincing interpretations of the results and conclusions. I highly recommend authors not to jump to interpreting emission/flux sources and patterns based on the single site measurements and the analysis done in this study. Sections in this manuscript dealing with these aspects need to be (preferably) removed or highly restructured.

While revising the manuscript, We have focused on the key aspect of the data: (i) better interpretation of the results, as suggested by the reviewer, and (ii) removed or refined speculative sentences and (iii) removed the whole section 3.5 (influence of regional emission and transport), which the reviewer found less convincing. Also we have substantially revised the interpretation of emission sources based on CO/CO_2 ratios, clearly stating that the results are only indicative, not confirmative (lines 715-730). In addition to the points noted above, we hope that by

addressing all the specific comments collectively we have adequately addressed the referee's general comments as well.

Specific comments and suggestions for revised analysis

Page 3, lines 63-66 "Between 1750…..cover changes" Misleading sentence. The given estimation is overly high for the accumulated CO2 in the atmosphere. As per IPCC reports & other studies, it is in the range of 230-250 Pg C. The given number is more towards total (cumulative) CO2 emissions between 1750 and 2011, which were partly compensated by the ocean and terrestrial ecosystems.

We would like to thank reviewer for noticing the misleading sentence which is now corrected with CO_2 accumulation data from the IPCC 2013 report as follows (lines 66-67):

Between 1750 and 2011, 240($\pm$10) PgC of anthropogenic CO_2 was accumulated in the atmosphere…. (IPCC, 2013).

Page 4, line 86-89 "important sources" Give reference

The following papers have been cited (see changed line 90-94) and the full citation is also included in the reference section.

Central Bureau of Statistics (CBS).: Nepal Living Standards Survey 2010/11, Statistical Report Volume 1, Central Bureau of Statistics, Government of Nepal, 2011.

Pandey, A., Sadavarte, P., Rao A. B., Venkataraman, C.: Trends in multipollutant emissions from a technology-linked inventory for India: II. Residential, agricultural and informal industry sectors, Atmos. Environ., 99, 341-352, doi: 10.1016/j.atmosenv.2014.09.080, 2014.

Sinha, V., Kumar, V., and Sarkar, C.: Chemical composition of pre-monsoon air in the Indo-Gangetic Plain measured using a new air quality facility and PTR-MS: high surface ozone and strong influence of biomass burning, Atmos. Chem. Phys., 14, 5921-5941, 10.5194/acp-14-5921-2014, 2014.

Page 4, lines 90-92 "Ecosystem and.... between July-October Please give appropriate reference for this statement. Prasad et al., 2014 investigated based on satellite GHG concentration observations, not based on inverse or ecosystem models. In my knowledge there's no such inverse flux estimations available over South Asia, accounting (and decoupling) seasonal variations of CO2 uptake and release to the atmosphere. However, Patra et al., 2011 showed inverse estimations of monthly co2 fluxes over South Asia. Please correct it.

Thank you for pointing out the incorrect statement. The sentences and reference have been replaced by relevant reference in lines 95-100 as follows:

By using inverse modeling, Patra et al. (2011) found a net CO_2 uptake $(0.37 \pm 0.20$ Pg C $yr^{-1})$ during 2008 in South Asia and the uptake (sink) is highest during July-September. The remaining months act as a weak gross sink but a moderate gross source for CO_2 in the region.

Page 12, lines 308-310 "2.193 (_ 0.224) ppm, 419.4 (_ 23.9) ppm, 0.50 (_ 0.35) ppm, and 310 1.71 (_ 0.71) %" Uncertainty seems to be quite high for co2, ch4 and co. Why? Please include the reason in the text.

We apologize for not clarifying about the values in the parenthesis adjacent to the annual average value for the three species. They are not measurement uncertainties. Instead, they are one standard deviation, calculated from the hourly data for the observation period (of one year). The reported annual mean from the references sites such as Mauna Loa and referred to in this paper were calculated from monthly means, not hourly data. Therefore for consistency and ease of comparison, we have now reported annual average and the standard deviation from monthly mean data for all sites. Please see the correction (lines 316-322):

For the entire sampling period, the annual average ($\pm$ one standard deviation) of CH_4, CO_2, CO, and water vapor mixing ratios at Bode were 2.192 ($\pm$ 0.066) ppm, 419.3 ($\pm$ 6.0) ppm, 0.50 ($\pm$ 0.23) ppm, and 1.73 ($\pm$ 0.66)% respectively. The relative variabilities for CH_4, CO_2 and CO were thus 3%, 1.4% and 46%, respectively. Their variabilities at Mauna Loa were CH4: 6% and CO_2: 0.5% and at Waligaun were CH_4: 0.48%, CO_2: 0.9%. The high variability in the annual mean,

notably for CO in Bode could be indicative of the seasonality of emission sources and meteorology.

Page 12, lines 318-320 "CH4 was ….. observation period" Given the above uncertainty ranges of Bode values (nearly 10.2% for CH4, 5.7% for CO2 and 70% for CO),the estimated percentages of increment relative to other observatories are also biased. These uncertainties need to be taken into account, or at least properly mentioned.

As noted above, these are not uncertainties, rather variabilities. We have recalculated the annual average and standard deviation values based on monthly average data. Now, the variabilities for CH_4, CO_2 and CO at Bode are 3%, 1.4% and 46%, respectively. The comparison between Bode and the other sites (Mauna Loa and Waligaun) for the GHGs now looks more reasonable. The following text has been inserted to reflect their variabilities and statistical significance (lines 330-334)

We performed a significance test at 95% confidence level (t-test) of the annual mean values between the sites to evaluate whether the observed difference is statistically significant ($p < 0.05$), which was confirmed for the annual mean CH_4 and CO_2 between Bode and Mauna Loa, and between Bode and Waliguan.

Page 12, lines 320-322 "The small .... Asia region" Although I tend to agree that we can expect (+ seeing satellite images), most of the cases, higher CH4 mixing ratios in Asia relative to Mauna Loa observatory, authors should note that this conclusion, as given in the text, about the whole Asia cannot be drawn from analysing "just two Asian sites" in a given time period. This sentence is misleading, and needs to be reformulated.

Thanks for drawing our attention to it. Yes, we didn't intend to generalize our comparison for Asia. We rephrased the following paragraph (lines 334-340):

CH_4 was nearly 20% higher at Bode than at Mauna Loa (1.831 ± 0.110 ppm) (Dlugokencky et al., 2017) and ca.17% higher than at Mt. Waliguan (1.879 ± 0.009 ppm) for the same observation

period (Dlugokencky et al., 2016). The slightly higher CH_4 mixing ratios at Bode and Waliguan than at Mauna Loa Observatory could be due to prevalence of rice farming as a key source of CH_4 in this part of Asia.

Page 12, line 324 "Ahmedabad (1.880 ppm) (Sahu and Lal, 2006) and Shadnagar (1.92 _ 0.07)" See my above comment on uncertainty.

We have calculated the variabilities for Bode, Ahmedabad and Shadnagar, and rephrased the sentences in lines 340-345 as follows:

Similarly, the annual average CH_4 at Bode during 2013-14 was found comparable to an urban site in Ahmedabad (1.880 ± 0.4 ppm, i.e., variability: 21.3%) in India for 2002 (Sahu and Lal, 2006) and 14% higher than in Shadnagar (1.92 ± 0.07 ppm, i.e., variability: 3.6%), a semi-urban site in Telangana state (~70 km north from Hyderabad city) during 2014 (Sreenivas et al., 2016).

Page 12, lines 326-332 "Likewise, the .... in China" These estimated increments are meaningless based on the uncertainty range of Bode's tracer mixing ratios. That is, these 5.7% and 5.5 % increments are statistically insignificant when it is compared with CO2 values with 5.7% uncertainty range. I strongly recommend authors to remove this.

See above regarding the confusion between uncertainty and variability, for which we apologize that this was not clear. We have recalculated the standard deviations based on the monthly values, and have also conducted a t-test between the annual mean mixing ratios and found that differences between the means at Bode and Mauna Loa, and between Bode and Waliguan were statistically significant. We have removed unnecessary explanations and inserted the following sentence (lines 345-348):

Likewise, the difference between annual mean CO_2 mixing ratios at Bode (419.2 ±6.0 ppm, 1.4% variability) vs. Mauna Loa (396.8 ± 2.0 ppm, 0.5% variability) (NOAA, 2015) and Bode vs. Waliguan (397.7 ± 3.6 ppm, 0.9% variability) (Dlugokencky et al., 2016a) is statistically significant ($p < 0.05$).

Page 13, lines 354-355 "burning activities due to rainfall in the region" What about the burning activities in Bode area during this rainy period of time?

Similar to Chanban, the burning activities around Bode area were also reduced or absent during the rainy season. We added the following line (378-379) in the manuscript.

The garbage and agro-residue burning activities were also absent or reduced around Bode due to rainfall during the monsoon period.

Page 14, lines 384-385 "The seasonal .... residue burn" What about seasonality of atmospheric transport and other met. fields? I would think that it can also make a significant impact especially considering mountain valley effect.

Agree. The paragraph has been restructured in lines 408-414 as:

The seasonal variation in CO_2 could be due to (i) the seasonality of major emission sources such as brick kilns, (ii) seasonal growth of vegetation (CO_2 sink) (Patra et al., 2011) and (iii) atmospheric transport associated with regional synoptic atmospheric circulation (monsoon circulation and westerly disturbance in spring season) which could transport regional emission sources from vegetation fires and agriculture residue burning (Putero et al., 2015), and a local mountain-valley circulation effect (Kitada and Regmi, 2003; Panday et al., 2009).

Page 14, lines 388 "partially due to rain washout. " It can also very well due to relatively high advection and vertical mixing.

Thank you for pointing out other possible reasons of low mixing ratios of gas species in rainy season. They are included (lines 414-422):

The concentrations of most pollutants in the region are lower during the monsoon period (Sharma et al., 2012, Marinoni, 2013; Putero et al., 2015) because frequent and heavy rainfall

suppresses emission sources. We saw a drop in the CO_2 mixing ratio during the rainfall period due to changes in various processes such as enhanced vertical mixing, uptake of CO2 by vegetation and soils, and where relevant reduction in combustion sources. CO_2 can also dissolve into rainfall, forming carbonic acid, which may lead to a small decrease in the CO_2 mixing ratio as has been observed during heavy intensity rainfall (Mahesh et al., 2014; Chaudhari et al., 2007).

Page 15, lines 395-396 "related to less or no rainfall, which results in the absence of rain washout" if it is with transport or raining, it should also affect CH4 and other tracers. Please clarify.

The additional source of CH_4 is due to agricultural activities from the paddy fields in the monsoon season (esp. August-September) which is absent from October. Therefore we see the drop in CH_4 starting October. It is likely that the absence of rainfall after October is conducive to CO_2 accumulation and thus we see an increase in the mixing ratio of CO_2 thereafter. We have tried to clarify it in lines 432-433 as:

However, the reduction in ambient CH_4 after October could be due to reduced CH_4 emissions from paddy fields, which were high in August-September.

Page 15, lines 412-413 "which had > 2.5 ppm CH4 and > 450 CO2" It's likely that the increase in CO2 & CH4 is associated with advected signals from the North East and the East; however it is not clear how it can be interpreted as the advected airmass had the said values for CO2 and CH4 unless the study used any tracer transport and emission models. Wind direction and tracer concentration from the given site alone are not sufficient to conclude this. A clarification is needed here. Otherwise I recommend authors to remove this. Page 15, lines 414 "(not shown in Figure 4)" I encourage authors to show CO as well for the completeness of the interpretation.

We agree that in the absence of tracer transport and emission modeling, it is not sufficient to conclude the directionality of the advected signals. The sentence formulation also gave an

impression that we are definitive about our conclusion. We have revised the lines 448-452. And CO has been included in Figure 4 as suggested by reviewer.

In the absence of tracer model simulations, the directionality of the advected air masses is unclear. Figure 4 shows that during these two months, CO_2 mixing ratios were particularly high (> 450 CO_2 and > 2.5 ppm CH_4) with the air masses coming from the Northeast-East (NE-E).

Page 15, lines 417 "high CH4 emissions" What about CO2?

The CO_2 level is lower during monsoon period in Asia due to high uptake of CO_2 by plants (high photosynthetic activities). This has been mentioned in lines 95-100 and 422-423 in the manuscript with reference.

âˇAˊlPage 17, line 455 "The westerly circulation (originated at longitude about 60E in 5 days back trajectories)" What does it mean? What's originated? Talking about model? Based on modeled trajectories?

The air mass back trajectory analysis is based on the study by Putero et al., 2015 for a site in the Kathmandu Valley for the period February 2013-January 2014. Using the Hysplit trajectory model, Putero et al (2015) clustered/grouped the 5 day back trajectories into a total of 9 clusters. The "westerly circulation" referred to in the manuscript is one of the dominant clusters (21.4%) observed by Putero. If spatially viewed on a lambert conformal projection, the majority of individual trajectories in the westerly circulation "or westerly cluster" originates in the region of $20-40^0$ N, $\sim 60^0$ E.

We acknowledge that the sentence was confusing. Please see the following correction which, we hope, provides a fairly detailed and clear explanation (lines 495-504).

To relate the influence of synoptic circulation with the observed variability in BC and O_3 in the Kathmandu Valley, 5-day back trajectories (of air masses arriving in the Kathmandu Valley) were computed by Putero et al., (2015) using the HYSPLIT model. These individual trajectories which were initialized at 600 hPa, for the study period of one year, and were clustered into nine

clusters. Of the identified clusters, the most frequently observed clusters during the study period were the Regional and Westerly cluster or circulation (22% and 21%). The trajectories in the regional cluster originate within 10° x 10° around the Kathmandu Valley, whereas the majority of trajectories in this westerly cluster originated broadly around 20-40^0 N, ~60^0 E. Putero et al (2015) found that the regional and westerly synoptic circulation were favorable for high values of BC and O_3 in the Kathmandu Valley.

Page 17, lines 475-476 "CO2 mixing ratios whereas CO shows an evening peak" It's surprising. Why is it so? Please clarify.

We meant to say that the CO_2 and CH_4 keep increasing over evening time until early morning while CO shows a peak coinciding with evening peak traffic hour and then drops. The decay in CO is more pronounced in monsoon and post-monsoon seasons. We have tried to clarify further in lines 526-534.

The gradual increase of CO_2 and CH_4 mixing ratios in the evening in contrast to the increase until evening peak traffic hours and later decay of CO may be indicative of a few factors. As pointed out earlier, after the peak traffic hours, there are no particularly strong sources of CO, especially in the monsoon and post-monsoon season. It is also likely that some of the CO decay is due to nighttime katabatic winds which replace polluted air masses with cold and fresh air from the nearby mountain (Panday and Prinn, 2009). As for the CO_2, the biosphere respiration at night in the absence of photosynthesis can add additional CO_2 to the atmosphere, which especially in the very shallow nocturnal boundary layer may explains part of the increase of the CO_2 mixing ratio.

Page 19, lines 519-521 "While the .... other seasons" I couldn't follow how it's related. A clarification is highly needed. What are these other most CO sources mentioned here?

We have tried to make it clear and also included other CO sources in lines 578-581. The new sentences read as:

While the biosphere respires at night, which may cause a notable increase in CO_2 in the shallow boundary layer, most CO sources (transport sector, residential cooking) except brick kilns remain shut down or less active at night.

Page 20, lines 551-559 "Highest day ....Mauna Loa andWaliguan" It's lost. I see many assumptions here rather than convincing statements. What about biospheric activity and its seasonality? mesoscale transport mechanisms?

We acknowledge that the whole paragraph was not clear. We have tried to make it concise and clear and also included some of your suggestions. The paragraph has been revised as follows (lines 612-628):

The highest daytime minimum of CO_2 was observed in the pre-monsoon, followed by winter (Figure 6b). The higher daytime minimum of CO_2 mixing ratios in the pre-monsoon season than in other seasons, especially winter, is interesting. The local emission sources are similar in pre-monsoon and winter and the boundary layer is higher (in the afternoon) during the pre-monsoon (~1200 meters) than in winter (~900 meters) (Mues et al., 2017). Also, the biospheric activity in the region is reported to be higher in the pre-monsoon (due to high temperature and solar radiation) than winter (Rodda et al., 2016). Among various possible causes, transport of CO_2 rich air from outside the Kathmandu Valley has been hypothesized as a main contributing factor, due to regional vegetation fire combined with westerly mesoscale to synoptic transport (Putero et al. 2015). In monsoon and post-monsoon seasons, the minimum CO_2 mixing ratio in the afternoon drops down to 390 ppm, this was close to the values observed at the regional background sites Mauna Loa and Waliguan.

Rodda, S.R., Thumaty, K. C., Jha, C. S. and Dadhwal, V. K.: Seasonal Variations of Carbon Dioxide, Water Vapor and Energy Fluxes in Tropical Indian Mangroves. Forests, 7, 35; doi:10.3390/f7020035, 2016.

Page 21, lines 577-579 "Overall, the .... fire etc." Importantly it shares the transport mechanisms.

We would like to thank reviewer for raising the role of transport mechanism in this section and it has been added in lines 647-650. The revised sentence reads as follows:

Overall, the positive and high correlations between CH_4 and CO mixing ratios and between CH_4 and CO_2 in the pre-monsoon and winter indicate common sources, most likely combustion related sources such as vehicular emission, brick kilns, agriculture fire etc., or the same source regions (i.e. their transport from outside the Kathmandu Valley due to regional atmospheric transport mechanisms).

Pages 21-22, Sec. 3.5 I strongly recommend authors to remove the whole section. It is not at all straight forward, as assumed here, to determine the impact of emission sources and transport, based on concentration measurements from single site. It does not make any sense unless a dedicated further study is involved to justify the stated assumptions here. The section, as in the present shape, does not meet scientific reasoning; hence need to be removed.

We are in agreement with the reviewer that without any good supporting evidence, it can be difficult to accept the assumptions made. Thus, as suggested, the whole section has removed from the manuscript and the remaining sections are renumbered.

Page 23, lines 633-641 "Based on the .... post-monsoon seasons" This could be a likely scenario. Have these interpretations been supported by any emission inventories available? It is also important to point out the associated uncertainties involved in separating different emission sectors based on this approach. Note that this approach cannot separate near and far field sources, different lifetimes of tracers etc.

Yes, we agree that the source identification based on CO/CO_2 ratio is indicative, not a definitive evidence. We apologize if we sounded definitive in our sentences and conclusion derived from the ratio analysis. We have also stated clearly in the paragraph that this method or the values associated with the source (Table 5) or our estimated values may have large uncertainty. A few additional lines have been added 680-682 to convey our cautious approach. Although we were unable to estimate the standard deviation of the ratio in Table 5, we have included the standard

deviation of our calculated ratio in Table 6, and added useful statistics such as geometeric mean and geometric standard deviation and their upper and lower bounds. The whole paragraph is significantly modified, and as supporting evidence, we also have added emission source sectors in the Kathmandu Valley based on a high resolution emission inventory (Sadavarte et al., 2017, in preparation).. Here is the modified paragraph (lines 715-730):

Although ratio of CO/CO_2 is a weak indicator of sources and the mean ratio has large variance (See Table 6), the conclusions drawn from using Figure 8 and the above mentioned classification are not conclusive. The estimated CO/CO_2 ratio tentatively indicates that the local plume impacting the measurement site (Bode) from the north and east could be residential and/or diesel combustion. The estimated CO/CO_2 ratio of the local plume from the south and west generally falls in the 15-45 range, which could indicate emissions from brick kilns and inefficient gasoline vehicles. Very high ratios were also estimated from the south west during the post-monsoon season. Among other possible sources, this may indicate agro-residue open burning.

The emission inventory for CO identifies (aggregate for a year) residential, and gasoline related emission from transport sector (Sadavarte et al., 2017, in preparation). The inventory is not yet temporally resolved, so no conclusion can be drawn about the sources with respect to different seasons. From the 1km x1km emission inventory of the Kathmandu Valley for 2011, the estimated sectoral source apportionment of CO is residential (37%), transport sector (40%) and industrial (20%). The largest fraction from the residential sector is cooking (24%) whereas the majority of transport sector related CO in the Kathmandu Valley is from gasoline vehicles.

Reference:

Sadavarte P., Das B., Shakya K., Rupakheti M., Bhave P. V., Byanju R. M., Lawrence M. G.: A High Resolution Technology-based Bottom-up Emissions Inventory for Nepal (manuscript in preparation), 2017.

Page 25, lines 704-706 "but it is clear that" What makes it clear?

We have removed the confusing words and rephrased the sentence in lines 793-795 as:

The cause of this diurnal pattern at Chanban is presently unclear, but the levels could be representative of the regional background throughout the day and show only limited influences of local emissions.

Page 26, lines 731-732 "Regional transport .... during pre-monsoon" I don't see any valid justification for this throughout.

The sentence has been removed from the conclusion section because we also realized that there is not a justifiable connection with the result presented in the manuscript.

Page 27, lines 743-744 "Low values .... mixing ratios." Please provide supporting details.

We have rephrased the sentence (lines 832-833):

Low values of CH_4 and CO_2 mixing ratios at the Chanban site could represent regional background mixing ratios.

Page 27, line 746 "useful for evaluation of satellite measurements and climate" How?

We have included the reason how the measurement is useful for satellite evaluation and rephrased the sentence to make it clear (lines 835-837):

These observations can be useful as ground-truthing for evaluation of satellite measurements, as well as climate and regional air quality models.

Page 27, lines 747-749 "The analysis ... Kathmandu Valley", Please, remove this sentence. Note that this is not met here and the study only demonstrates the observational variations of GHGs in the study region.

We agree with the reviewer that this study focuses on observational variations of mainly CH_4 and CO_2. However, this study along with recent studies will help in addressing mitigation of the pollution in the study region. Thus, we have rephrased the sentence in lines 837-839as:

The overall analysis presented in the paper will contribute along with other recent measurements and analysis to providing a sound scientific basis for reducing emissions of greenhouse gases and air pollutants in the Kathmandu Valley.

Reference Please check. Formatting issues and sometimes journal details (or other important parts) are missing

We have re-checked the references and tried to resolve formatting and other issues in the section.

Table 4 Columns 4 & 5: What are these values? Looks like monthly values for Bode during Aug., and Sep. 2013; but then why are they different from corresponding 2nd column of the table 4?

These are monthly values for Bode, which were simultaneously measured with another site at Chanban during August and September 2015. The information has been included in Table 4.

Table 6 Column 2: This "*" meant for?

We would like to thank reviewer for noticing this symbol. The meaning of the symbol has been included with a note at the bottom of the Table 6 as:

*The morning peak was one hour delayed in winter, thus the 8:00-10:00 period data was used in the analysis.

Figure 4 I didn't follow the fig. well. What I understood is that the plot shows the frequency of hourly mixing ratios w.r.t the frequency of prevailing wind direction. Did it also take into account the wind speed? Then what about different percentages shown? For example, does it

mean that 5% of sample time in August, the wind was from NE and "CH4_corrected" is above 2.5 ppm in which less than 1% time (in my eyes), CH4 is in 3-3.3 ppm range? In that case, it's statistically difficult to say that the monthly enhancement is due to the polluted air masses from the NE and E. By the way, what are these "_corrected" values for CH4 and CO2? What about March-April scenario for CO2_corrected? Did two plots (Figs.2 and 3) use same set of master data, or any quality filtering had been done other than monthly averaging?

We are sorry for creating confusion by keeping "CH4_corrected" in the legend of Figure 4. CO_2 and CH_4 corrected means water corrected values of them, which we explained in section 2.2….. The CH_4 and CO_2 data used in the whole analysis (including Figure 2 and 3) are the same and they are water corrected values. We did not use any additional filtering while making these plots in the analysis. To avoid confusion, we have replaced CH_4_corrected and CO_2_corrected by CH_4 and CO_2 respectively in the legend of Figure 4.

1   **Seasonal and diurnal variations of methane and carbon dioxide in the Kathmandu Valley**

2   **in the foothills of the central Himalaya**

3   Khadak Singh Mahata[1,2], Arnico Kumar Panday [3,4], Maheswar Rupakheti[1,5*], Ashish Singh[1],

4   Manish Naja[6], Mark G. Lawrence[1,2]

5   [1] Institute for Advanced Sustainability Studies (IASS), Potsdam, Germany

6   [2] University of Potsdam, Potsdam, Germany

7   [3] International Centre for Integrated Mountain Development (ICIMOD), Lalitpur, Nepal

8   [4] University of Virginia, Virginia, USA

9   [5] Himalayan Sustainability Institute (HIMSI), Kathmandu, Nepal

10   [6] Aryabhatta Research Institute of Observational Sciences (ARIES), Nainital, India

11

12   *Correspondence to: M. Rupakheti (maheswar.rupakheti@iass-potsdam.de)

13

14   **Abstract**

15   The SusKat-ABC (Sustainable Atmosphere for the Kathmandu Valley- Atmospheric Brown

16   Clouds) international air pollution measurement campaign was carried out during December

17   2012-June 2013 in the Kathmandu Valley and surrounding regions in Nepal. The Kathmandu

18   Valley is a bowl-shaped basin with a severe air pollution problem. This paper reports

19   measurements of two major greenhouse gases (GHGs), methane (CH_4) and carbon dioxide

20   (CO_2), along with the pollutant CO, that began during the campaign and were extended for a

21   year at the SusKat-ABC supersite in Bode, a semi-urban location in the Kathmandu Valley.

22   Simultaneous measurements were also made during 2015 in Bode and a nearby rural site

23   (Chanban), ~25 km (aerial distance) to the southwest of Bode, on the other side of a tall ridge.

24   The ambient mixing ratios of methane (CH_4), carbon dioxide (CO_2), water vapor, and carbon

25   monoxide (CO) were measured with a cavity ring down spectrometer (Picarro G2401, USA),

along with meteorological parameters for a year (March 2013 - March 2014). Simultaneous measurements were also made at Chanban from 15 July to 3 October 2015. These measurements are the first of their kind in the central Himalayan foothills. At Bode, the annual average mixing ratios of CO_2 and CH_4 were 419.34(±236.09) ppm and 2.1923(±0.066224) ppm, respectively. These values are higher than the levels observed at background sites such as Mauna Loa, USA (CO_2: 396.8 ± 2.0 ppm, CH_4: 1.831 ± 0.110 ppm) and Waliguan, China (CO_2: 397.7 ± 3.6 ppm, CH_4: 1.879 ± 0.009 ppm) during the same period, and at other urban/semi-urban sites in the region such as Ahmedabad and Shadnagar (India) and Nanjing (China). They varied slightly across the seasons at Bode, with seasonal average CH_4 mixing ratios being 2.157(±0.230) ppm in the pre-monsoon season, 2.199(±0.241) ppm in the monsoon, 2.210(±0.200) ppm in the post-monsoon, and 2.214(± 0.209) ppm in the winter season. The average CO_2 mixing ratios were 426.2(±25.5) ppm in pre-monsoon, 413.5(±24.2) ppm in monsoon, 417.3(±23.1) ppm in post-monsoon, and 421.9(±20.3) ppm in winter season. The maximum seasonal mean mixing ratio of CH_4 in winter was only 0.057 ppm or 2.6% higher than the seasonal minimum during the pre-monsoon period, while CO_2 was 12.8 ppm or 3.1% higher during the pre-monsoon period (seasonal maximum) than during the monsoon (seasonal minimum). On the other hand, the CO mixing ratio at Bode was 191% higher during the winter than during the monsoon season. The enhancement in CO_2 mixing ratios during the pre-monsoon season is associated with additional CO_2 emissions from forest fire and agro-residue burning in northern South Asia in addition to local emissions in the Kathmandu Valley. Published CO/CO_2 ratios of different emission sources in Nepal and India were compared with the observed CO/CO_2 ratios in this study. This comparison indicated that the major sources in the Kathmandu Valley were residential cooking and vehicle exhaust in all seasons except winter. In winter, the brick kiln emissions were a major source. Simultaneous measurement in Bode and Chanban (15 July-3 Oct 2015) revealed that the mixing ratio of CO_2, CH_4 and CO mixing ratios were 3.8%, 12%, and 64% higher in Bode than Chanban. Kathmandu Valley, thus, has significant emissions from local sources, which can also be attributed to its bowl shaped geography that is conducive to pollution build-up. At Bode, Aall three gas species (CO_2, CH_4 and CO) in Bode showed strong diurnal patterns in their mixing ratios with a pronounced morning peak (ca. 08:00), a dip in the afternoon, and again gradual

55 increase through the night until the next morning, whereas CH_4 and CO at Chanban did not show

56 any noticeable diurnal variations.

57 These measurements provide the first insights into diurnal and seasonal variation of key

58 greenhouse gases and air pollutants and their local and regional sources, which are important

59 information for the atmospheric research in the region.

1   Introduction

61 The average atmospheric mixing ratios of two major greenhouse gases (GHGs), CO_2 and CH_4,

62 have increased by about 40% (from 278 to 390.5 ppm) and about 150% (from 722 to 1803 ppb)

63 respectively since pre-industrial times (~1750 AD). This is mostly attributed to anthropogenic

64 emissions (IPCC, 2013). The current global annual rate of increase of the atmospheric CO_2

65 mixing ratio is 1-3 ppm, with average annual mixing ratios now exceeding a value of 400 ppm at

66 the background reference location in Mauna Loa (WMO, 2016). Between 1750 and 2011,

67 240555(±1085) PgC of anthropogenic CO_2 was accumulatedadded to in the atmosphere of which

68 two thirds were contributed by fossil fuel combustion and cement production, with the remaining

69 coming from deforestation and land use/land cover changes (IPCC, 2013). CH_4 is the second

70 largest gaseous contributor to anthropogenic radiative forcing after CO_2 (Forster et al., 2007).

71 The major anthropogenic sources of atmospheric CH_4 are rice paddies, ruminants and fossil fuel

72 use, contributing approximately 60% to the global CH_4 budget (Chen and Prinn, 2006;

73 Schneising et al., 2009). The remaining fraction is contributed by biogenic sources such as

74 wetlands and fermentation of organic matter by microbes in anaerobic conditions (Conrad,

75 1996).

76 Increasing atmospheric mixing ratios of CO_2 and CH_4 and other GHGs and short-lived climate-

77 forcing pollutants (SLCPs) such as black carbon (BC) and tropospheric ozone (O_3) have caused

78 the global mean surface temperature to increase by 0.85°C from 1880 to 2012. The surface

79 temperature is expected to increase further by up to 2 degrees at the end of the 21st century in

80 most representative concentration pathways (RCP) emission scenarios (IPCC, 2013). The

81 increase in surface temperature is linked to melting of glaciers and ice sheets, sea level rise,

82 extreme weather events, loss of biodiversity, reduced crop productivity, and economic losses

83 (Fowler and Hennessy, 1995; Guoxin and Shibasaki, 2003).

84 Seventy percent of global anthropogenic CO_2 is emitted in urban areas (Fragkias et al., 2013).

85 Developing countries may have lower per capita GHG emissions than developed countries, but

86 the large cities in developing countries, with their high population and industrial densities, are

87 major consumers of fossil fuels and thus, emitters of GHGs. South Asia, a highly populated

88 region with rapid growth in urbanization, motorization, and industrialization in recent decades,

89 has an ever increasing fossil fuel demand and its combustion emitted 444 Tg C/year in 2000

90 (Patra, et al., 2013), or about 5% of the global total CO_2 emissions. Furthermore, a major

91 segment of the population in South Asia has an agrarian economy and uses biofuel for cooking

92 activities, and agro-residue burning is also common practice in the region, which areis an

93 important major sources of air pollutants and greenhouse gases in the region (CBS, 2011; Pandey

94 et al., 2014; Sinha et al., 2014).

95 The emission and uptake of CO_2 and CH_4 follow a distinct cycle in South Asia. By using inverse

96 modeling, Patra et al. (2011) found also showed thata net CO_2 uptake (0.37 ± 0.20 Pg C yr^{-1})

97 during 2008 in South Asia and the uptake (sink) is highest peaks induring July-September. And

98 tThe remaining months acts as a weak gross sink but a moderate gross source for CO_2 in the

99 regionAugust, using an inversion constrained by regional measurements from commercial

100 aircraft. The observed trend variation is linked with the growing seasons. Agriculture is a major

101 contributor of methane emission. For instance, in India it contributes to 75% of CH_4 emissions

102 (MoEF, 2007). Ambient CH_4 concentrations are highest during June to September (peaking in

103 September) in South Asia which are also the growing months for rice paddies (Goroshi et al.,

104 2011). The minimum column averaged ambient CH_4 concentrations mixing ratios are observed

105 in February-March (Prasad et al., 2014).

106 Climate change has impacted South Asia in several ways, as evident in temperature increase,

107 change in precipitation patterns, higher incidence of extreme weather events (floods, droughts,

108 heat waves, cold waves), melting of snowfields and glaciers in the mountain regions, and

109 impacts on ecosystems and livelihoods (ICIMOD, 2009; MoE, 2011). Countries such as Nepal

110 are vulnerable to impacts of climate change due to inadequate preparedness for adaptation to
111 impacts of climate change (MoE, 2011). Decarbonization of its economy can be an important
112 policy measure in mitigating climate change. Kathmandu Valley is one of the largest
113 metropolitan cities in the foothills of the Hindu Kush-Himalaya which has significant reliance on
114 fossil fuels and biofuels. In 2005, fossil fuel burning accounted for 53% of total energy
115 consumption in the Kathmandu Valley, while biomass and hydroelectricity were 38% and 9%,
116 respectively (Shrestha and Rajbhandari, 2010). Fossil fuel consumed in the Kathmandu Valley
117 accounts for 32% of the country's fossil fuel imports, and the major fossil fuel consumers are
118 residential (53.17%), transport (20.80%), industrial (16.84%), and commercial (9.11%) sectors.
119 Combustion of these fuels in traditional technologies such as Fixed Chimney Bulls Trench Kiln
120 (FCBTK) and low efficiency engines (vehicles, captive power generator sets etc.) emit
121 significant amounts of greenhouse gases and air pollutants. This has contributed to elevated
122 ambient concentrations of particulate matter (PM), including black carbon and organic carbon,
123 and several gaseous species such as ozone, polycyclic aromatic hydrocarbons (PAHs),
124 acetonitrile, benzene and isocyanic acid (Pudasainee et al., 2006; Aryal et al., 2009; Panday and
125 Prinn, 2009; Sharma et al., 2012; World Bank, 2014; Chen et al., 2015; Putero et al., 2015:
126 Sarkar et al., 2016). The ambient levels often exceed national air quality guidelines (Pudasainee
127 et al., 2006; Aryal et al., 2009; Putero et al., 2015) and are comparable or higher than ambient
128 levels observed in other major cities in South Asia.

129 Past studies in the Kathmandu Valley have focused mainly on a few aerosols species (BC, PM)
130 and short-lived gaseous pollutants such as ozone and carbon monoxide (Pudasainee et al., 2006;
131 Aryal et al., 2009; Panday and Prinn, 2009; Sharma et al., 2012, Putero et al., 2015). To the best
132 of authors' knowledge, no direct measurements of CO_2 and CH_4 are available for the Kathmandu
133 Valley. Recently, emission estimates of CO_2 and CH_4 were derived for the Kathmandu Valley
134 using the International Vehicle Emission (IVE) model (Shrestha et al., 2013). The study
135 estimated 1554 Gg of annual emission of CO_2 from a fleet of vehicles (that consisted of public
136 buses, 3-wheelers, taxis and motor cycles; private cars, trucks and non-road vehicles were not
137 included in the study) for the year 2010. In addition, the study also estimated 1.261 Gg of CH_4
138 emitted from 3 wheelers (10.6 %), taxis (17.7 %) and motorcycles (71 %) for 2010.

27

139   This study presents the first 12 months of measurements of two key GHGs, CH_4 and CO_2 along

140   with other trace gases and meteorological parameters in Bode, a semi-urban site in the eastern

141   part of the Kathmandu Valley. The year-long measurement in Bode is a part of the SusKat-ABC

142   (Sustainable Atmosphere for the Kathmandu Valley – Atmospheric Brown Clouds) international

143   air pollution measurement campaign conducted in and around the Kathmandu Valley from

144   December 2012 to June 2013. Details of the SusKat-ABC campaign are described in Rupakheti

145   et al. (2017<s>6</s>, manuscript in preparation). The present study provides a detailed account of

146   seasonal and diurnal behaviors of CO_2 and CH_4 and their possible sources. To examine the rural-

147   urban differences and estimate the urban enhancement, these gaseous species were also

148   simultaneously measured for about three months (Jul-Oct) in 2015 at Chanban, a rural site about

149   25 km (aerial distance) outside and southwest of Kathmandu Valley. The seasonality of the trace

150   gases and influence of potential sources in various (wind) directions are further explored by via

151   ratio analysis. This measurement provides unique data from highly polluted but relatively poorly

152   studied region (central Himalayan foothills in South Asia) which could be useful for validation

153   of emissions estimates, model outputs and satellite observations. The study, which provides new

154   insights on potential sources, can also be a good basis for designing mitigation measures for

155   reducing emissions of air pollutants and controlling greenhouse gases in the Kathmandu Valley

156   and the region.

157   2   Experiment and Methodology

158   2.1 Kathmandu Valley

159   The Kathmandu Valley consists of three administrative districts: Kathmandu, Lalitpur, and

160   Bhaktapur, situated between 27.625° N, 27.75° N and 85.25°E, 85.375°E. It is a nearly circular

161   bowl-shaped valley with a valley floor area of approximately 340 km^2 located at an altitude of

162   1300 m mean sea level (masl). The surrounding mountains are close to 2000-2800 in height

163   above sea level with five mountain passes located at about 200-600 m above the valley floor and

164   an outlet for the Bagmati River southwest of the Kathmandu Valley. Lack of decentralization in

165   in Nepal has resulted in the concentration of economic activities, health and education facilities,

166   the service sector, as well as most of the central governmental offices in the Kathmandu Valley.

167   Consequently, it is one of the fastest growing metropolitan areas in South Asia with a current
168   population of about 2.5 million, and the population growth rate of 4% per year (World Bank,
169   2013) Likewise, approximately 50% of the total vehicle fleet (2.33 million) of the country is in
170   Kathmandu Valley (DoTM, 2015). The consumption of fossil fuels such as liquefied petroleum
171   gas (LPG), kerosene for cooking and heating dominates the residential consumption, while the
172   rest use biofuel (fuelwood, agro-residue, animal dung) for cooking and heating in the Kathmandu
173   Valley. The commercial sector is also growing in the valley, and the latest data indicate the
174   presence of 633 industries of various sizes. These are mainly associated with dyeing, brick kilns,
175   and manufacturing industries. Fossil fuels such as coal and biofuels are the major fuels used in
176   brick kilns. Brick kilns are reported as one of the major contributors of air pollution in the
177   Kathmandu Valley (Chen et al., 2015; Kim et al., 2015; Sarkar et al., 2016). There are about 115
178   brick industries in the valley (personal communication with M. Chitrakar, President of the
179   Federation of Nepalese Brick Industries). Acute power shortage in the Valley is common all
180   around the year, especially in the dry season (winter/pre-monsoon) when the power cuts can last
181   up to 12 hours a day (NEA, 2014). Energy demand during the power cut period is met with the
182   use of small (67% of 776 generators surveyed for the World Bank study was with capacity less
183   than 50kVA) but numerous captive power generators (diesel/petrol), which further contribute to
184   valley's poor air quality. According to the World Bank's estimate, over 250,000 such generator
185   sets are used in the Kathmandu Valley alone, producing nearly 200 MW of captive power, and
186   providing about 28% of the total electricity consumption of the valley (World Bank, 2014).
187   Apart from these sources, trash burning, which is a common practice (more prevalent in winter)
188   throughout the valley, is one of the major sources of air pollutants and GHGs.

189   Climatologically, Kathmandu Valley has a sub-tropical climate with annual mean temperature of
190   18°C, and annual average rainfall of 1400 mm, of which 90% occurs in monsoon season (June-
191   September). The rest of the year is dry with some sporadic rain events. The wind circulation at
192   large scale in the region is governed by the Asian monsoon circulation and hence the seasons are
193   also classified based on such large scale circulations and precipitation: Pre-Monsoon (March-
194   May), Monsoon (June-September), Post-Monsoon (October-November) and Winter (December-
195   February). Sharma et al. (2012) used the same classification of seasons while explaining the

196 seasonal variation of BC concentrations observed in the Kathmandu Valley. Locally in the
197 valley, the mountain-valley wind circulations play an important role in influencing air quality.
198 The wind speed at the valley floor is calm (≤ 1 m s^{-1}) in the morning and night, while a westerly
199 wind develops after 11:00 AM in the morning till dusk, and switches to a mild easterly at night
200 (Panday and Prinn, 2009; Regmi et al., 2003). This is highly conducive to building up of air
201 pollution in the valley, which gets worse during the dry season.

2.2 Study sites

203 Two sites, a semi-urban site within the Kathmandu Valley and a rural site outside the Kathmandu
204 Valley, were selected for this study. The details of the measurements carried out in these sites is
205 described Table 1 and in section 2.2.1 and 2.2.2.

2.2.1 Bode (SusKat-ABC supersite)

207 The SusKat-ABC supersite was set up at Bode, a semi-urban location (Figure 1) of the
208 Madhyapur Thimi municipality in the Bhaktapur district in the eastern side of the Kathmandu
209 Valley. The site is located at 27.68^0N latitude, 85.38^0E longitude, and 1344 masl. The local area
210 around the site has a number of scattered houses and agricultural fields. The agriculture fields are
211 used for growing rice paddies in the monsoon season. It also receives outflow of polluted air
212 from three major cities in the valley: Kathmandu Metropolitan City and Lalitpur Sub-
213 metropolitan City, both mainly during daytime, and Bhaktapur Sub-metropolitan City mainly
214 during nighttime. Among other local sources around the site, about 10 brick kilns are located in
215 the east and southeast direction, approximately within 1-4 km from the site which are operational
216 only during dry season (January to April). There are close to 20 small and medium industries
217 (pharmaceuticals, plastics, electronics, tin, wood, aluminum, iron, and fabrics etc.) scattered in
218 the same direction. The Tribhuvan International Airport (TIA) is located approximately 4 km
219 away to the west of the Bode site.

2.2.2 Chanban

221 Chanban is a rural/background site in Makwanpur district outside of the Kathmandu Valley
222 (Figure 1). This site is located ~25 km aerial distance due southwest from Bode. The site is

223 located on a small ridge (27.65^0N, 85.14^0E, 1896 masl) between two villages - Chitlang and

224 Bajrabarahi - within the forested watershed area of Kulekhani Reservoir, which is located ~ 4.5

225 km southwest of the site. The instruments were set up on the roof of 1-storey building in an open

226 space inside the Nepali Army barrack. There was a kitchen of the army barrack at about 100 m to

227 the southeast of the measurement site. The kitchen uses LPG, electricity, kerosene, and firewood

228 for cooking activities.

2.3 Instrumentation

230 The measurements were carried out in two phases in 2013-2014 and 2015. In phase one, a cavity

231 ring down spectrometer (Picarro G2401, USA) was deployed in Bode to measure ambient CO_2,

232 CH_4, CO, and water vapor mixing ratios. Twelve months (6 March 2013 - 5 March 2014) of

233 continuous measurements were made in Bode. The operational details of the instruments

234 deployed in Bode are also provided in Table 1. In phase two, simultaneous measurements were

235 made in Bode and Chanban for a little less than 3 months (15 July to 03 October 2015).

236 The Picarro G2401 analyzer quantifies spectral features of gas phase molecules by using a novel

237 wavelength-scanned cavity ring down spectroscopic technique (CRDS). The instrument has a 30

238 km path length in a compact cavity that results high precision and sensitivity. Because of the

239 high precision wavelength monitor, it uses absolute spectral position and maintains accurate peak

240 quantification. Further, it only monitors the special features of interest for reducing the drift. The

241 instrument also has water correction to report dry gas fraction. The reported measurement

242 precision for CO_2, CH_4, CO and water vapor in dry gas is < 150 ppb, < 30 ppb, < 1ppb and < 200

243 ppm for 5 seconds with 1 standard deviation (Picarro, 2015).

244 In Bode, the Picarro analyzer was placed on the 4th floor of a 5-storey building with an inlet at

245 0.5 m above the roof of the building with a 360 degree view (total inlet height: 20 m above

246 ground). The sample air was filtered at the inlet to keep dust and insects out and was drawn into

247 the instrument through a 9 m Teflon tube (1/4 inches ID). The Picarro analyzer was set to record

248 data in every 5 second and recorded both directly sampled data and water corrected data of CO_2

249 and CH_4. In this paper, only water-corrected or dry mixing ratios of CH_4 and CO_2 were used to

250 calculate the hourly averages for diurnal and seasonal analysis.

251   The instruments were factory calibrated before commencing the field measurement. Picarro
252   G2401 model is designed for remote application and long term deployment with minimal drift
253   and less requirement for intensive calibration (Crosson, 2008) and thus was chosen for the
254   current study in places like Kathmandu where there is no or limited availability of high quality
255   reference gases. Regular calibration of Picarro G2401 in field during 2013-2014 deployment was
256   not conducted due to challenges associated with the quality of the reference gas, especially for
257   CO and CH_4. One time calibration was performed for CO_2 (at 395, and 895 ppmv) in July 2015
258   before commencing the simultaneous measurement in Bode and Chanban in 2015. The
259   %difference between CO_2 mixing ratio reported by of the analyzer and the differed
260   byapproximately 5% at reference mixing ratio was within 5%. CO observations from Picarro
261   G2401 were compared with observations from another CO analyzer (Horiba, model AP370) that
262   was also operated in Bode for 3 months (March - May 2013). Horiba CO monitor was a new
263   unit, which was factory calibrated before its first deployment in Bode. Nevertheless, this
264   instrument was inter-compared with another CO analyzer (same model) from the same
265   manufacturer prior to the campaign and its correlation coefficient was 0.9 [slope of data from the
266   new unit (y-axis) vs the old unit (x-axis) = 1.09]. Primary gas cylinders from Linde UK (1150
267   ppbv) and secondary gases from Ultra-Pure Gases and Chemotron Science Laboratories (1790
268   ppbv) were used for the calibration of CO instrument. Further details on CO measurements and
269   calibration of Horiba AP370 can be found in Sarangi et al. (2014; 2016). Statistically significant
270   correlation (r = 0.99, slope = 0.96) was found between Picarro and Horiba hourly average CO
271   mixing ratio data (Supplementary Information Figure S1). Furthermore, the monthly mean
272   difference between these two instruments (Horiba AP370 minus Picarro G2401) was calculated
273   to be 0.02 ppm (3%), 0.04 ppm (5%) and 0.02 ppm (4%) in March, April and May, respectively.
274   For the comparison period of 3 months, the mean difference was 0.02 ppm (4%). Overall
275   differences were small to negligible during the comparison period and thus, adjustment in the
276   data was deemed not necessary.

277   Besides highly selective to individual species, Picarro G2401 has a water correction function and
278   thus accounts for the any likely drift in CO, CO_2 and CH_4 mixing ratios with the fluctuating
279   water vapor concentration (Chen et al., 2013;Crosson, 2008). Crosson (2008) also estimated a

280   peak to peak drift of 0.25 ppmv. Further, Crosson (2008) observed a 1.2 ppbv/day drift in CO_2

281   after 170 days from the initial calibration. For a duration of one year the drift will be less than 1

282   ppmv, which is less than 1% of the observed mixing ratio in (hourly ranges: 376-537 ppm) Bode

283   even if the drift was in same magnitude as in case of Crosson (2008). Crosson (2008) reported

284   0.8 ppbv peak to peak drift in CH_4 measurements for 18 days after the initial calibration.

285   There were other instruments concurrently operated in Bode; a ceilometer for measuring mixing

286   layer height (Vaisala Ceilometer CL31, Finland), and an Automatic Weather Station (AWS)

287   (Campbell Scientific, USA). The ceilometer was installed on the rooftop (20 m above ground) of

288   the building (Mues et al., 2017). For measuring the meteorological parameters, a Campbell

289   Scientific AWS (USA) was set up on the roof of the building with sensors mounted at 2.9 m

290   above the surface of the roof (22.9 m from the ground). The Campbell Scientific AWS measured

291   wind speed and direction, temperature, relative humidity and solar radiation every minute.

292   Temperature and rainfall data were taken from an AWS operated by the Department of

293   Hydrology and Meteorology (DHM), Nepal at the Tribhuvan International Airport (TIA, see

294   Figure 1), ~4 km due west of Bode site.

295   At Chanban, the inlet for Picarro gas analyzer was kept on the rooftop ~3 m above the ground

296   and the sample air was drawn through a 3 m long Teflon tube (1/4 inches ID). The sample was

297   filtered at the inlet with a filter (5-6 µm pore size) to prevent aerosol particles from getting into

298   the analyzer. An automatic weather station (Davis Vantage Pro2, USA) was also set up in an

299   open area, about 17 m away from the building and with the sensors mounted at 2 m above

300   ground.

301   **3. Results and discussion**

302   The results and discussions are organized as follow: Sub-section 3.1 describes a year round

303   variation in CH_4, CO_2, CO and water vapor at Bode; sub-sections 3.2, 3.3 ~~and 3.4~~ present the

304   analysis of the observed ~~diurnal,~~ monthly and seasonal variations and diurnal variation. Sub-

305   sections 3.4 and 3.5~~, 3.6, 3.7~~ discusse~~s~~ ~~theimpact of city pollution at the measurement site at~~

306   ~~Bode, influence of regional pollution and transport,~~about the interrelation of CO_2, CH_4 and CO

307 and potential emission sources in the valley and sub-section 3.6~~8~~ compares and contrasts CH$_4$,

308 CO$_2$, CO at Bode and Chanban.

3.1 Time series of CH$_4$, CO$_2$, CO and water vapor mixing ratios

310 Figure 2 shows the time series of hourly mixing ratios of CH$_4$, CO$_2$, CO, and water vapor at

311 Bode. Meteorological data from Bode and the Tribhuvan International Airport are also shown in

312 Figure 2. Data gaps in Figure 2a and 2b were due to maintenance of the measurement station. In

313 general, the changes observed in CO mixing ratio was higher in terms of % change than the

314 variations observed in CH$_4$ and CO$_2$ mixing ratios during the sampling period. In contrast, CO

315 mixing ratios decreased and water vapor mixing ratios increased significantly during the rainy

316 season (June-September). For the entire sampling period, the annual average (± one standard

317 deviation) of CH$_4$, CO$_2$, CO, and water vapor mixing ratios were 2.19~~23~~ (±0.066 ~~0.224~~) ppm,

318 419.3~~4~~ (±6.0 ~~23.9~~) ppm, 0.50 (±0.23 ~~0.35~~) ppm, and 1.73~~1~~ (±0.66 ~~0.71~~)%, respectively. The

319 relative variabilities for the annual average of CH$_4$, CO$_2$ and CO were thus 3%, 1.4% and 46%,

320 respectively. Their variabilities at Mauna Loa were CH$_4$: 6% and CO$_2$: 0.5% and at Waliguan

321 were CH$_4$: 0.48%, CO$_2$: 0.9%. The high variability in the annual mean, notably for CO in Bode

322 could be indicative of the seasonality of emission sources and meteorology. The annual CH$_4$ and

323 CO$_2$ mixing ratios were compared to the historical background site (Mauna Loa Observatory,

324 Hawaii, USA) and the background site (Waliguan, China) in Asia, which will provide insight on

325 spatial differences. The selection of neighboring ~~countries' (i.e.,~~ Indian ~~and China's)~~ urban and

326 semi-urban sites, where many emission sources are typical for the region, for comparison

327 provides information on relative differences (higher/lower), which will help in investigating

328 possible local emission sources in the valley. As expected, annual mean of CH$_4$ and CO$_2$ mixing

329 ratios in the Kathmandu Valley were higher than the levels observed at background sites in the

330 region and elsewhere ~~for the same period~~ (Table 4). We performed a significance test at 95%

331 confidence level (t-test) of the annual mean values between the sites to evaluate whether the

332 observed difference is statistically significant (p < 0.05), which was confirmed for the. ~~We found~~

333 ~~the difference in~~ annual mean CH$_4$ and CO$_2$ between Bode and Mauna Loa, and between Bode

334 and Waliguan ~~were statistically significant~~. CH$_4$ was nearly 2~~3~~0% higher at Bode than at Mauna

335 Loa observatory (1.831 ± 0.110 ppm) (Dlugokencky et al., 2017~~6~~) and ca.17% higher than at Mt.

34

336 Waliguan (1.879 ± 0.009 ppm) ~~in China~~ for the same observation period (Dlugokencky et al.,

337 2016). The slightly higher~~small difference~~ CH_4 mixing ratios between at Bode and Waliguan ~~in~~

338 ~~comparison to~~than at Mauna Loa Observatory could ~~signal~~indicates ~~the higher mixing ratio of~~

339 ~~CH_4 in these two Asian sites. The~~be due to rice farming as a key source of CH_4 in this part of

340 Asia. Thus, ~~I~~it could be associated with such agricultural activities in this region. Similarly, the

341 annual average CH_4 at Bode during 2013-14 was found comparable to ~~higher than urban/semi-~~

342 ~~urban sites in India, such as~~ an urban site in Ahmedabad (1.880 ± 0.4 ppm, i.e., variability:

343 21.3%) in India for 2002 (Sahu and Lal, 2006) and 14% higher than in Shadnagar (1.92 ± 0.07

344 ppm, i.e., variability: 3.6%), a semi-urban site in Telangana state (~70 km north from Hyderabad

345 city) during 2014 (Sreenivas et al., 2016). Likewise, the difference between annual mean mixing

346 ratios at Bode (419.3 ±6.0 ppm, 1.4% variability) vs. Mauna Loa (396.8 ± 2.0 ppm, 0.5%

347 variability) (NOAA, 2015) and Bode vs. Waliguan (397.7 ± 3.6 ppm, 0.9% variability)

348 (Dlugokencky et al., 2016a) is statistically significant (p <0.05). ~~the annual average CO_2 mixing~~

349 ~~ratio at Bode (419.4 ±23.9 ppm) during the observation period was 5.7% higher than at Mauna~~

350 ~~Loa Observatory (396.76 ppm) (Tans and Keeling, 2014) and 5.5% higher than at Mt. Waliguan~~

351 ~~(397.7 ppm). The CO_2 mixing ratio in the Kathmandu Valley was also found to be higher than~~

352 ~~the levels observed in Shadnagar (394 ± 2.9 ppm) during 2014, Ahmedabad city (413 ± 13.7~~

353 ~~ppm) in India during November 2013 to May 2015, and an urban site at Nanjing (406.5 ± 20~~

354 ~~ppm) in China (Huang et. al., 2015; Sreenivas et al., 2016; Chandra et al., 2016).~~

355

356 The high CH_4 and CO_2 mixing ratios at Bode in comparison to Ahmedabad and Shadnagar ~~and~~

357 ~~Nanjing~~ could be due to more than 115 coal-biomass fired brick kiln, some of them are located

358 near the site (less than 4 km) and confinement of pollutants within the Valley due to bowl shaped

359 topography of the Kathmandu Valley. Although Ahmedabad ~~and Nanjing sites are~~ is ~~in~~ a big

360 cit~~yies~~ with high population larger than Kathmandu Valley but the measurement site~~y is~~are far

361 from the nearby heavy polluting industries and situated in plains, where ventilation of pollutants

362 would be more efficient as opposed to the Kathmandu Valley. The major polluting sources were

363 industries, residential cooking and transport sector in Ahmedabad (Chandra et al., 2016).

364 ~~Anthropogenic emission, synoptic circulation, terrestrial biosphere had important role on CO_2~~

365 ~~mixing ratios in Nanjing (Huang et al., 2015)~~. Shadnagar is a small town with a population of

366 0.16 million and major sources were industries (small-medium), biomass burning in residential

367 cooking (Sreenivas et al., 2016).

368 The monthly average of CO_2 mixing ratios in 2015 in Chanban (Aug: 403.4, Sep: 399.1 ppm)

369 were slightly higher than the background sites at Mauna Loa Observatory (Aug: 398.89 ppm,

370 Sep: 397.63 ppm) (NOAA, 2015) and Mt. Waliguan (Aug: 394.55 ppm, Sep: 397.68 ppm)

371 (Dlugokencky et al., 2016a). For these two months in 2015, CH_4 mixing ratios were also higher

372 in Bode (Aug: 2.28~~1.11~~ ppm~~b~~, Sep: 2.37~~10.93~~ ppm~~b~~) and Chanban (Aug: 2.050~~49.71~~ ppm~~b~~,

373 Sep: 2.102~~1.75~~ ppm~~b~~) compared to Mauna Loa Observatory (Aug: 1.83~~1.04~~ ppm~~b~~, Sep:

374 1.846~~5.68~~ ppm~~b~~) (Dlugokencky et al., 2017~~6~~)) and Mt. Waliguan (Aug: 1.915~~4.99~~ ppm~~b~~,

375 1.911~~.21~~ ppm~~b~~) (Dlugokencky et al., 2016). The low differences in CO_2 between Chanban and

376 background sites mentioned above indicate the less number of and/or less intense CO_2 sources at

377 Chanban during these months because of the lack of burning activities due to rainfall in the

378 region. The garbage and agro-residue burning activities were also absent or reduced around Bode

379 during the monsoon period. However, high CH_4 values in August and September in Bode,

380 Chanban and Mt. Waliguan in comparison to Mauna Loa Observatory may indicate the influence

381 of CH_4 emission from paddy fields in the Asian region.

382 **3.2 Monthly and Seasonal variations**

383 Figure 3 shows the monthly box plot of hourly CH_4, CO_2, CO and water vapor observed for a

384 year in Bode. Monthly and seasonal averages of CH_4 and CO_2 mixing ratios at Bode are

385 summarized in Table 2 and 3. CH_4 were lowest during May-July (ranges from 2.093-2.129 ppm)

386 period and highest during August-September (2.274-2.301 ppm), followed by winter. In addition

387 to the influence of active local sources, the shallow boundary layer in winter was linked to

388 elevated concentrations (Panday and Prinn, 2009; Putero et al., 2015, Mues et al., 2017~~6~~). The

389 low CH_4 values from May to July may be associated with the absence of brick kiln and frequent

390 rainfall in these months. Brick kiln were operational during January to April. Rainfall also leads

391 to suppression of open burning activities in the valley (see Figure 2b). The CH_4 was slightly

392 higher (statistically significant, $p<0.05$) in monsoon season (July –September) than pre-monsoon

393 season (unlike CO_2 which was higher in pre-monsoon), and could be associated with the addition

394   of CH_4 flux from the water-logged rice paddies (Goroshi et al., (2011). There was a visible drop

395   in CH_4 from September to October but remained consistently over 2.183 ppm from October to

396   April with little variation between these months. Rice-growing activities are minimal or none in

397   October and beyond, and thus may be related to the observed dip in CH_4 mixing ratio.

398   Comparison of seasonal average CH_4 mixing ratios at Bode and Shadnagar (a semi-urban site in

399   India) indicated that CH_4 mixing ratios at Bode were higher in all seasons than at Shadnagar:

400   pre-monsoon (1.89 ± 0.05 ppm), monsoon (1.85 ± 0.03 ppm), post-monsoon (2.02 ± 0.01 ppm),

401   and winter (1.93 ± 0.05 ppm) (Sreenivas et al., 2016). The possible reason for lower CH_4 at

402   Shadnagar in all season could be associated with geographical location and difference in local

403   emission sources. The highest CH_4 mixing ratio in Shadnagar was reported in post-monsoon

404   which was associated with harvesting in the Kharif season (July – October), while the minimum

405   was in monsoon. Shadnagar is a relatively small city (population: ~0.16 million) compared to

406   Kathmandu Valley and the major local sources which may have influence on CH_4 emission

407   include bio-fuel, agro-residue burning and residential cooking.

408   The seasonal variation in CO_2 could be due to (i)generally reflects the seasonality of major

409   emission sources such as brick kilns and regional emission sources such as vegetation fire and

410   agriculture residue burn, (ii) seasonal growth of vegetation (CO_2 sink) (Patra et al., 2011) and

411   (iii) atmospheric transport associated with regional synoptic atmospheric circulation (monsoon

412   circulation and westerly disturbance in spring season) which could transport regional emission

413   sources from vegetation fire and agriculture residue burning (Putero et al., 2015), and a local

414   mountain-valley circulation effect (Kitada and Regmi, 2003; Panday et al., 2009). The

415   concentrations of most pollutants in the region are lower during the monsoon period (Sharma et

416   al., 2012, Marinoni, 2013; Putero et al., 2015) due to limited because frequent and heavy rainfall

417   suppresses emission sources and partially due to rain washout. We saw a drop in the CO_2 mixing

418   ratio during the rainfall period due to changes in various processes such as enhanced vertical

419   mixing, uptake of CO_2 by vegetation and soils, and where relevant reduction in combustion

420   sources. CO_2 can also dissolve into rainfall, forming carbonic acid, which may lead to a small

421   decrease in the CO_2 mixing ratio as has been observed during heavy intensity rainfall (Mahesh et

422   al., 2014; Chaudhari et al., 2007). Monsoon is also the growing season with higher CO_2

423   assimilation by plants than other seasons (Sreenivas et al., 2016). In contrast, winter, pre-
424   monsoon and post-monsoon season experiences an increase in emission activities in the
425   Kathmandu Valley (Putero et al., 2015).

426

427   The CO_2 mixing ratios were in the range of 376 - 537 ppm for the entire observation period.
428   Differences with CH_4 were observed in September and October where CO_2 was increasing
429   (mean/median) in contrast to CH_4 which showed the opposite trend. The observed increase in
430   CO_2 after October may be related to less or no rainfall, which results in the absence of rain-
431   washout and/or no suppression of active emission sources such as open burning activities.
432   However, the reduction in CH_4 after October could be due to reduced CH_4 emissions from paddy
433   fields, which were high in August-September. CO_2 remains relatively lower during July-August,
434   but it is over 420 ppm from January to May. Seasonal variation of CO_2 in Bode was similar in
435   seasonal variation but the values are higher than the values observed in Shadnagar, India
436   (Sreenivas et al., 2016).

437   The variations in CO were more distinct than CH_4 and CO_2 during the observation period
438   (Figure 3). The highest CO values were observed from January-April (0.71-0.91 ppm). The
439   seasonal mean of CO mixing ratios at Bode were: pre-monsoon (0.60 ±0.36 ppm), monsoon
440   (0.26±0.09 ppm), post-monsoon (0.40±0.15 ppm), and winter (0.76±0.43 ppm). The maximum
441   CO was observed in winter, unlike CO_2 which was maximum in pre-monsoon. The high CO in
442   winter was due to the presence of strong local pollution sources (Putero et al., 2015) and shallow
443   mixing layer heights. The addition of regional forest-fire and agro-residue burning augmented
444   CO_2 mixing ratios in pre-monsoon. The water vapor mixing ratio showed a seasonal pattern
445   opposite of CO, with a maximum in monsoon (2.53 %) and minimum in winter (0.95 %), and
446   intermediate values of 1.56 % in pre-monsoon and 1.55 % in post-monsoon season.

447   There were days in August-September when the CH_4 increases by more than 3 ppm (Figure 2).
448   Enhancement in CO_2 was also observed during the same time period. In the absence of tracer
449   model simulations, the directionality of the advected air masses is unclear. Figure 4 showsed that
450   during these two months, It is likely that these high enhancements CO_2 mixing rations were
451   particularly high (> 450 CO_2 and > 2.5 ppm CH_4) with the air masses coming in-from the were

452 ~~associated with the air mass from~~ Northeast-East (NE-E). ~~which had > 2.5 ppm CH₄ and > 450~~

453 ~~CO₂ (see Figure 4)~~ CO during the same period was not enhanced and didn't show any particular

454 directionality compared to CH_4 and CO_2 (~~not shown in~~ Figure 4c). Areas NE-E to Bode are

455 predominantly irrigated (rice paddies) during August-September, and sources such as brick kilns

456 were not operational during this time period. Goroshi et al. (2011) reported that June to

457 September is a growing season for rice paddies in South Asia with high CH_4 emissions during

458 these months and observed a peak in September in the atmospheric CH_4 column over India.

459 Model analysis also points to high methane emissions in September which coincides with the

460 growing period of rice paddies (Goroshi et al., 2011, Prasad et al., 2014). The CH_4 mixing ratios

461 at Bode in January (2.233 ± 0.219 ppm) and July (2.129 ± 0.168 ppm) were slightly higher than

462 the observation in Darjeeling (Jan: 1.929±0.056 ppm; Jul: 1.924±0.065 ppm), a hill station of

463 eastern Himalaya (Ganesan et al., 2013). The higher CH_4 values in January and July at Bode

464 compared to Darjeeling could be because of the influence of local sources, in addition to the

465 shallow boundary layer in Kathmandu Valley. Trash burning and brick kilns are two major

466 sources from December until April in the Kathmandu Valley while emission from paddy fields

467 occurs during July-September in the Kathmandu Valley. In contrast, the measurement site in

468 Darjeeling was located at higher altitude (2194 masl) and was less influenced by the local

469 emission. The measurement in Darjeeling reflected a regional contribution. There are limited

470 local source in Darjeeling such as wood biomass burning, natural gas related emission and

471 vehicular emission (Ganesan et al., 2013).

472 The period between January and April had generally higher or the highest values of CO_2, CH_4

473 and CO at Bode. The measurement site was impacted mainly by local Westerly-Southwesterly

474 winds (W-SW) and East-Southeast (E-SE). The W-SW typically has a wind speed in the range

475 ~1 - 6 m s^{-1} and was active during late morning to afternoon period (~11:00 to 17:00 NST,

476 supplementary information Figure S2 and S3). Major cities in the valley such as Kathmandu

477 Metropolitan City and Lalitpur Sub-metropolitan City are W-SW of Bode (Figure 1c). Wind

478 from E-SE were generally calm (≤1m s^{-1}) and observed only during night and early morning

479 hours (21:00 to 8:00 NST). The mixing ratio of all three species in air mass from the E-SE was

480 significantly higher than in the air mass from W-SW (Figure 4). There are 10 biomass co-fired

481 brick kilns and Bhaktapur Industrial Estate located within 1-4 km E-SE from Bode (Sarkar et al.,

482 2016). The brick kilns were only operational during January-April. Moreover, there were over

483 100 brick kilns operational in the Kathmandu Valley (Putero et al., 2015) which use low-grade

484 lignite coal imported from India and biomass fuel to fire bricks in inefficient kilns (Brun, 2013).

485 Fresh emissions from main city center were transported to Bode during daytime by W-SW winds

486 which mainly include vehicular emission. Compared to monsoon months (June-August), air

487 mass from W-SW had higher values in all three species (Figure 4) during winter and pre-

488 monsoon months. This may imply that in addition to vehicular emission, there are other potential

489 sources which were exclusively active during these dry months. Municipal trash burning is also

490 common in the Kathmandu Valley, with a reported higher frequency from December to February

491 (Putero et al., 2015). The frequency in the use of captive power generator sets are highest during

492 the same period, which is another potential source contributing to air coming from W-SW

493 direction (World Bank, 2014; Putero et al., 2015).

494 Regional transport of pollutants into the Kathmandu Valley was reported by Putero et al. (2015).

495 To relate the influence of synoptic circulation with the observed variability in BC and O_3 in the

496 Kathmandu Valley, 5-day back trajectories (of air masses arriving in the Kathmandu Valley)

497 were computed by Putero et al., (2015) using the HYSPLIT model. These individual trajectories

498 which were initialized at 600 hPa, for the study period of one year and were clustered into nine

499 clusters. Of the identified clusters, the most frequently observed clusters during the study period

500 were the Regional and Westerly cluster or circulation (22% and 21%). The trajectories in the

501 regional cluster originate within 10° x 10° around the Kathmandu Valley, whereas the majority of

502 trajectories in this westerly cluster originated broadly around 20-40^0 N, ~60^0 E. Putero et al

503 (2015) found that ~~T~~the regional and westerly synoptic circulation were favorable for high values

504 of BC and O_3 in the Kathmandu Valley. ~~(originated at longitude about 60^{0}E in 5 days back~~

505 ~~trajectories) was dominant from March May 2013.~~ Other sources of CO_2 and CH_4 could be due

506 to vegetation fires which were also reported in the region surrounding the Kathmandu Valley

507 during the pre-monsoon months (Putero et al., 2015). Similarly, high pollution events, peaking in

508 the pre-monsoon, were observed at Nepal Climate Observatory-Pyramid (NCO-P) near Mt.

509 Everest, which have been associated with vegetation fires in the Himalayan foothills and

40

510 northern IGP region (Putero et al., 2014). MODIS derived forest counts (Figure 5), which also

511 indicated high frequency of forest fire and farm fire from February to April and also during post-

512 monsoon season. It is interesting that the monthly mean CO_2 mixing ratio was maximum in April

513 (430 ± 27 ppm) which could be linked to the fire events. It is likely that the westerly winds

514 (>2.5-4.5 m s^{-1}) during the daytime (supplementary information Figure S2, S3) bring additional

515 CO_2 from vegetation fires and agro-residue burning in southern plains of Nepal including the

516 IGP region (Figure 5). Low values of CO_2 and CH_4 during June-July (Figure 3) was coincident

517 with the rainy season, and sources such as brick kiln emission, trash burning, captive power

518 generators, and regional agriculture residue burning and forest fires are weak or absent during

519 these months.

3.3 Diurnal Variation

521 Figure 6 shows the average seasonal diurnal patterns of CH_4, CO_2, CO, and water vapor mixing

522 ratios observed at Bode for four seasons. All the three gas species had a distinct diurnal pattern in

523 all seasons, characterized by maximum values in the morning hours (peaked around 7:00-9:00),

524 afternoon minima around 15:00-16:00, and a gradual increase through the evening until next

525 morning. There was no clear evening peak in CH_4 and CO_2 mixing ratios whereas CO shows an

526 evening peak around 20:00. The gradual increase of CO_2 and CH_4 in the evening in contrast to

527 the increase until evening peak traffic hours and later decay of CO may be indicative of a few

528 factors. As pointed out earlier, after the peak traffic hours, there are no particularly strong

529 sources of CO, especially in the monsoon and post-monsoon season. It is also likely that some of

530 the CO is decay due to nighttime katabatic winds which replace polluted air masses with cold

531 and fresh air from the nearby mountain (Panday and Prinn, 2009). As for the CO_2, the biosphere

532 respiration at night in the absence of photosynthesis can add additional CO_2 to the atmosphere

533 which especially in the very shallow nocturnal boundary layer may explains part of the increase

534 of the CO_2 mixing ratio. The well-defined morning and evening peaks observed in CO mixing

535 ratios are associated with the peaks in traffic and residential activities. The CH_4 and CO_2 showed

536 pronounced peaks in the morning hours (07:00-09:00) in all seasons with almost the same level

537 of seasonal average mixing ratios. CO had a prominent morning peak in winter and pre-monsoon

538 season, but the peak was significantly lower in monsoon and post-monsoon. The CO (~1-1.4

41

ppm) around 08:00-09:00 am in winter and pre-monsoon were nearly 3-4 times higher than in monsoon and post-monsoon season. It appears that CH_4 and CO_2 mixing ratios were continuously building up at night until the following morning peak in all seasons. The similar seasonal variations in CH_4 and CO_2 across all seasons could be due to their long-lived nature, as compared to CO, whose diurnal variations are strongly controlled by the evolution of the boundary layer. Kumar et al. (2015) also reported morning and evening peaks and an afternoon low in CO_2 mixing ratios in industrial, commercial, and residential sites in Chennai in India. The authors also found high early morning CO_2 mixing ratios at all sites and attributed it to the temperature inversion and stable atmospheric condition.

The daytime low CH_4 and CO_2 mixing ratios were due to (i) elevated mixing layer height in the afternoon (Figure 7), (ii) development of upslope wind circulation in the valley, and (iii) development of westerly and southwesterly winds which blows through the valley during the daytime from around 11 am to 5 pm (supplementary information Figure S2), all of which aid in dilution and ventilation of the pollutants out of the valley (Regmi et al., 2003; Kitada and Regmi, 2003; Panday and Prinn, 2009). In addition, the daytime CO_2 minimum in the summer monsoon is also associated with high photosynthetic activities in the valley as well as in the broader surrounding region. In the nighttime and early morning, the mixing layer height was low (only around 200-300 m in all seasons) and stable boundary layer for almost 17 hours a day. In the daytime it grows up to 800-1200 m for a short time (ca. from 11:00 to 6:00) (Mues et al., 20176, manuscript submitted to ACPD). Therefore the emissions from various activities in the evening after 18:00 (cooking and heating, vehicles, trash burning, and bricks factories in the night and morning) were trapped within the collapsing and shallow boundary layer, and hence mixing ratios were high during evening, night and morning hours. Furthermore, plant and soil respiration also increases CO_2 mixing ratio during the night (Chandra et al., 2016). However, Ganesan et al. (2013) found a distinct diurnal cycle of CH_4 mixing ratios with twin peaks in the morning (7:00-9:00), and afternoon (15:00-17:00) and a nighttime low in winter but no significant diurnal cycle in the summer of 2012 in Darjeeling, a hill station (2194 masl) in the eastern Himalaya. The authors described that the morning peaks could be due to the radiative heating of the ground in the morning, which breaks the inversion layer formed during night, and as a result, pollutants are

569   ventilated from the foothills up to the site. The late afternoon peaks match wind direction and
570   wind speed (upslope winds) that could bring pollution from plains to mountains.

571   The diurnal variation of CO is also presented along with CO_2 and CH_4 in Figure 6c. CO is an
572   indicator of primary air pollution. Although CO mixing ratio showed distinct diurnal pattern, it
573   was different from the diurnal patterns of CO_2 and CH_4. CO diurnal variation showed distinct
574   morning and evening peaks, afternoon minima, and a nighttime accumulation or decay.
575   Nighttime accumulation in CO was observed only in winter and pre-monsoon and decay or
576   decrease in monsoon season and post-monsoon season (Figure 7). The lifetime of CO (weeks to
577   months) is very long compared to the ventilation timescales for the valley, so the different
578   diurnal cycles would be due to differences in nighttime emissions. While the biosphere respires
579   at night which may cause a notable increase in CO_2 in the shallow boundary layer, most CO
580   sources (transport sector, residential cooking) except brick kilns remain shut down or less active
581   late at night. This also explains why nighttime values of CO drop less in the winter and pre-
582   monsoon than in other seasons. Furthermore, the prominent morning peaks of CO in pre-
583   monsoon and winter compared to other seasons results from nighttime accumulation, additional
584   fresh emissions in the morning and recirculation of the pollutants due to downslope katabatic
585   winds (Pandey and Prinn, 2009; Panday et al., 2009). Pandey and Prinn (2009) observed
586   nighttime accumulation and gradual decay during the winter (January 2005). The measurement
587   site in Pandey and Prinn (2009) was near the urban core of the Kathmandu Valley and had
588   significant influence from the vehicular sources all over the season including the winter season.
589   Measurement in Bode lies in close proximity to the brick kilns which operate 24 hours during the
590   winter and pre-monsoon period. Calm southeasterly winds are observed during the nighttime and
591   early morning (ca.22:00 – 8:00) in pre-monsoon and winter, which transport emissions from
592   brick kiln to the site (Sarkar et al., 2016). Thus the gradual decay in CO was not observed in
593   Bode.

594   The timing of the CO morning peak observed in this study matches with observations by Panday
595   et al. (2009). They also found CO morning peak at 8:00 in October 2004 and at 9:00 in January
596   2005. The difference could be linked to the boundary layer stability. As the sun rises later in

597 winter, the boundary layer stays stable for a longer time in winter keeping mixing ratios higher in
598 morning hours than in other seasons with an earlier sunrise.

599 The morning peaks of CO_2 and CH_4 mixing ratios occurred around 6:00-7:00 local time in the
600 pre-monsoon, monsoon, and post monsoon season, whereas in winter their peaks are delayed by
601 1-2 hours in the morning; CH_4 at 8:00 and CO_2 at 9:00. The CO showed that its morning peak
602 was delayed compared to CO_2 and CH_4 morning peaks by 1-2 hour in pre-monsoon, monsoon
603 and post-monsoon (at 8:00) and in winter (at 9:00). The occurrence of morning peaks in CO_2 and
604 CH_4 1-2 hours earlier than CO is interesting. This could be due to the long lifetimes and
605 relatively smaller local sources of CH_4 and CO_2, as CO is mainly influenced by emissions from
606 vehicles during rush hour, as well as from biomass and trash burning in the morning hours. Also,
607 CO increases irrespective of change in mixing layer (collapsing or/rising, Figure 7) but CO_2 and
608 CH_4 start decreasing only after the mixing layer height starts to rise. Recently, Chandra et al.
609 (2016) also reported that the CO_2 morning peak occurred earlier than CO in observations in
610 Ahmedabad City India. This was attributed to CO_2 uptake by photosynthetic activities after
611 sunrise but CO kept increasing due to emissions from the rush hour activities.

612 The Hhighest daytime minimum of CO_2 was observed in the pre-monsoon followed by winter
613 (Figure 6b). which may indicate the influence of regional emissions that increased the baseline
614 background concentrations as wellThe higher daytime minimum of CO_2 mixing ratios in the pre-
615 monsoon season than in other seasons, especially winter, is interesting. The local emission
616 sources are similar in pre-monsoon and winter and the boundary layer is higher (in the afternoon)
617 during the pre-monsoon (~1200 meters) than in winter (~900 meters) (Mues et al., 2017). occurs
618 from 12:00 to 17:00 LST. The highest minimum CO_2 was found in pre-monsoon (Figure 6b).
619 Although the local emission sources are similar in pre-monsoon and winter Also, the biospheric
620 activity in the region is reported to be higher in the pre-monsoon (due to high temperature and
621 solar radiation) than winter (Rodda et al., 2016).Among various possible causes, transport of
622 CO_2 rich air from outside the Kathmandu Valley has been hypothesized as a main contributing
623 factor, due to regional vegetation fire combined with westerly mesoscale to synoptic transport ,
624 the higher minimum daytime CO_2 mixing ratios in pre-monsoon season than other seasons,
625 suggest the influence of regional emissions in the Kathmandu Valley, which has been reported in

44

626 ~~previous study by~~ Putero et al. (2015). In monsoon and post-monsoon seasons, the minimum

627 CO_2 mixing ratio~~s~~ in the afternoon drops down to 390 ppm, ~~which~~ this wa~~s~~ere close to the value~~s~~

628 observed at the regional background sites ~~such as~~ Mauna Loa and Waliguan.

629

3.4 Seasonal interrelation of CO_2, CH_4 and CO

631 The Pearson's correlation coefficient (r) between CO_2 and CO was strong in winter (0.87),

632 followed by monsoon (0.64), pre-monsoon (0.52) and post-monsoon (0.32). The higher

633 coefficient in winter indicates that common or similar sources for CO_2 and CO and moderate

634 values in pre-monsoon and monsoon indicates the likelihood of different sources. To avoid the

635 influence of strong diurnal variations observed in the valley, daily averages, instead of hourly,

636 were used to calculate the correlation coefficients. The correlation coefficients between daily

637 CH_4 and CO_2 for four seasons are as follows: winter (0.80), post-monsoon (0.74), pre-monsoon

638 (0.70) and monsoon (0.22). A semi-urban measurement study in India also found a strong

639 positive correlation between CO_2 and CH_4 in the pre-monsoon (0.80), monsoon (0.61), post-

640 monsoon (0.72) and winter (0.8) (Sreenivas et al., 2016). It should be noted here that Sreenivas

641 et al., (2006) used hourly average CO_2 and CH_4 mixing ratios. The weak monsoon correlation at

642 Bode, which is in contrast to Sreenivas et al. (2016), may point to the influence of dominant CH_4

643 emission from paddy field during the monsoon season (Goroshi et al., 2011). Daily CH_4 and CO

644 was also weakly correlated in monsoon (0.34) and post-monsoon (0.45). Similar to CH_4 and

645 CO_2, the correlation between CH_4 and CO were moderate to strong in pre-monsoon (0.76) and

646 winter (0.75).

647 Overall, the positive and high correlations between CH_4 and CO mixing ratios and between CH_4

648 and CO_2 in the pre-monsoon and winter indicate common sources ~~or source regions~~, most likely

649 combustion related sources such as vehicular emission, brick kilns, agriculture fire etc., or the

650 same source regions (i.e. their transport due to regional atmospheric transport mechanisms).

651 Weak correlation, between CH_4-CO_2 and between CH_4-CO, during monsoon season indicates

652 sources other than combustion-related may be active, such as agriculture as a key CH_4 source

653 (Goroshi et al., 2013)

3.6 3.5 CO and CO_2 ratio: Potential emission sources

The ratio of the ambient mixing ratios of CO and CO_2 was used as an indicator to help discriminate emission sources in the Kathmandu Valley. The ratio was calculated from the excess (dCO and dCO_2) relative to the background values of ambient CO and CO_2 mixing ratios. The excess value was estimated by subtracting the base value which was calculated as the fifth percentile of the hourly data for a day (Chandra et al., 2016).

Average emission ratios from the literature are shown in Table 5, and average ratios of dCO/dCO_2 are shown in Table 6, disaggregated into morning hours, evening hours, and seasonal values. It must be stated that due to the large variance in the calculated ratio from this study

(Table 6) as well as the likely variation in the estimated ratio presented in Table 5, the interpretation and conclusion about sources should be cautiously drawn and will be indicative. Higher ratios were found in pre-monsoon (12.4) and winter (15.1) season compared to post-monsoon (8.3) and monsoon (7.5). These seasonal differences in the dCO/dCO_2 ratio are depicted in Figure 8, which shows a clear relationship with the wind direction and associated emissions, with the highest values especially for stronger westerly winds. Compared to the other three seasons, the ratio in winter was also relatively high for air masses from the east, likely due to emissions from brick kilns combined with accumulation during more stagnant meteorological conditions (supplementary information Figure S2, S3). In other seasons, emission emanating from the north and east of Bode were characterized by a dCO/dCO_2 ratio below 15. Air masses from the west and south generally have a ratio from 20 to 50 in all but post-monsoon season, where the ratio sometimes exceeds 50. A ratio of 50 or over is normally due to very inefficient combustion sources (Westerdahl et al., 2009; Stockwell et al., 2016), such as agro-residue burning, which is common during the post-monsoon season in the Kathmandu Valley.

For interpretability of emission ratio with sources, the ratio was classified into three categories: (i) 0 – 15, (ii) 15 – 45, and (iii) greater than 45. This classification was based on the observed distribution of emission ratio during the study period (Figure 8) and a compilation of observed emission ratios typical for different sources from Nepal and India (see Table 5). An emission ratio below 15 is likely to indicate residential cooking and diesel vehicles, and captive power generation with diesel-powered generator sets (Smith et al., 2000; ARAI, 2008; World Bank, 2014). The emission from brick kilns (FCBTK and Clamp kilns, both common in the Kathmandu Valley), and inefficient, older (built before 2000) gasoline cars fall in between 15 - 45 (Weyant et al., 2014, Stockwell et al., 2016; ARAI, 2008). Four-stroke motorbikes and biomass burning activities (mixed garbage, crop-residue and biomass) are one of the least efficient combustion sources, with emission ratios higher than 45 (Westerdahl et al., 2009; Stockwell et al., 2016; ARAI, 2008).

Based on the classification and Figure 8, the emissions from sources to the north and east of the site are dominated by residential cooking and/or diesel combustion. Emissions from the south and west of Bode are mainly contributed by sources such as brick kilns and inefficient gasoline

710 ~~vehicles. Very high ratios, indicative of agro-residue open burning, generally only show up~~
711 ~~during the post-monsoon period, when such activities take place, especially in areas southwest of~~
712 ~~the site. The relatively enhanced ratio (20-30) observed in winds from north and east of the site~~
713 ~~during winter is mostly likely due to brick kilns that use mixed coal-biomass fuel, whereas the~~
714 ~~Figure 8 indicates the dominant signature of residential cooking, diesel and old gasoline cars~~
715 ~~during the pre-monsoon, monsoon and post-monsoon seasons.~~ Although ratio of CO/CO_2 is a
716 weak indicator of sources and the mean ratio has large variance (See Table 6), the conclusions
717 drawn, from using Figure 8 and the above mentioned classification, are not conclusive. The
718 estimated CO/CO_2 ratio tentatively indicates that the local plume impacting the measurement site
719 (Bode) from the north and east could be residential and/or diesel combustion. The estimated
720 CO/CO_2 ratio of the local plume from the south and west generally falls in the 15-45 range
721 which could indicate emissions from brick kilns and inefficient gasoline vehicles. Very high
722 ratios were also estimated from the south west during the post-monsoon season. Among other
723 possible sources, this may indicate agro-residue open burning.

724 The emission inventory for CO identifies (aggregate for a year) residential, and gasoline related
725 emission from transport sector (Sadavarte et al., 2017, in preparation*). The inventory is not yet
726 temporally resolved, so no conclusion can be drawn about the sources with respect to different
727 seasons. From the 1km x1km emission inventory of the Kathmandu Valley for 2011, the
728 estimated sectoral source apportionment of CO is residential (37%), transport sector (40%) and
729 industrial (20%). The largest fraction from the residential sector is cooking (24 %) whereas the
730 majority of transport sector related CO in the Kathmandu Valley is from gasoline vehicles.

731 The dCO/dCO_2 ratio also changes markedly between the morning peak hours (7:00-9:00, except
732 in winter season when the peak occurs during 8:00-9:00) and evening peak hours (19:00-21:00
733 pm) (Table 6). Morning and evening values were lowest (2.2, 8.0) during the monsoon and
734 highest (11.2, 21.6) in the winter season, which points to the different emission characteristics in
735 these two seasons. This feature is similar to Ahmedabad, India, another urban site in south Asia,
736 where the morning/evening values were lowest (0.9/19.5) in monsoon and highest in winter
737 (14.3/47.2) (Chandra et al., 2016). In the morning period, the ratio generally falls within a
738 narrower range, from less than 1 to about 25, which indicates a few dominant sources, such as

739 cooking, diesel vehicles, and diesel gen-sets (see Figure 9). In the evening period, the range of
740 the ratio is much wider, from less than 1 to more than 100, especially in winter. This is partly due
741 to the shallower boundary layer in winter, giving local CO emissions a chance to build up more
742 rapidly compared to the longer-lived and well-mixed CO_2, and also indicating the prevalence of
743 additional sources such as brick kilns and agro-residue burning.

744 **3.6~~7~~ Comparison of CH_4 and CO_2 at semi-urban site (Bode) and rural site (Chanban)**

745 Figure 10 shows time series of hourly average mixing ratios of CH_4, CO_2, CO and water vapor
746 observed simultaneously at Bode and Chanban for the period of 15th July to 3rd October 2015.
747 The hourly meteorological parameters observed at Chanban are shown in supplementary Figure
748 S4. The hourly temperature ranges from 14 to 28.5 °C during the observation period. The site
749 experienced calm winds during the night and moderate southeasterly winds with hourly
750 maximum speed of up to 7.5 m s^{-1} during the observation period. The CH_4 mixing ratios at
751 Chanban varied from 1.880 ppm to 2.384 ppm, and generally increased from the last week of
752 July until early September, peaking around 11th September and then falling off towards the end
753 of the month. CO followed a generally similar pattern, with daily average values ranging from
754 0.10 ppm to 0.28 ppm. The hourly CO_2 mixing ratios ranged from 375 to 453 ppm, with day to
755 day variations, but there were no clear pattern as observed in trend like CH_4 and CO mixing
756 ratios.

757 The CH_4, CO_2, and CO mixing ratios were higher in Bode than in Chanban (Figure 10, Table 4),
758 with Chanban approximately representing the baseline of the lower envelope of the Bode levels.
759 The mean CO_2, CH_4 and CO mixing ratios over the entire sampling period of nearly three
760 months at Bode are 3.8%, 12.1%, and 64% higher, respectively, than at Chanban. The difference
761 in the CO_2 mixing ratio could be due to the large uptake of CO_2 in the forested area at Chanban
762 and surrounding regions compared to Bode, where the local anthropogenic emissions rate is
763 higher and less vegetation for photosynthesis. The coincidence between the base values of CO
764 and CH_4 mixing ratios at Bode and the levels observed at Chanban implies that Chanban CO and
765 CH_4 mixing ratios are indicative of the regional background levels. A similar increase in CO and
766 CH_4 mixing ratios at Chanban from July to September was also observed at Bode, which may

49

767 imply that the regional/background levels in the broader Himalayan foothill region also
768 influences the baseline of the daily variability of the pollutants in the Kathmandu Valley,
769 consistent with Panday and Prinn (2009).

770 Figure 11 shows the comparison of average diurnal cycles of CO_2, CH_4, CO and water vapor
771 mixing ratios observed at Bode and Chanban. The diurnal pattern of CO_2 mixing ratios at both
772 sites is similar, but more pronounced at Bode, with a morning peak around 6:00-7:00, a daytime
773 minimum, and a gradual increase in the evening until the next morning peak. A prominent
774 morning peak at Bode during the monsoon season indicates the influence of local emission
775 sources. The daytime CO_2 mixing ratios are also higher at Bode than at Chanban because of local
776 emissions less uptake of CO_2 for photosynthesis in the valley in comparison to the forested area
777 around Chanban. Like the diurnal pattern of CO_2 depends on the evolution of the mixing layer at
778 Bode, as discussed earlier, it is expected that the mixing layer evolution similarly influences the
779 diurnal CO_2 mixing ratios at Chanban. CO, on the other hand, shows very different diurnal
780 patterns at Bode and Chanban. Sharp morning and evening peaks of CO are seen at Bode,
781 indicating the strong local polluting sources, especially cooking and traffic in the morning and
782 evening peak hours. Chanban, in contrast, only has a subtle morning peak and no evening peak.
783 After the morning peak, CO sharply decreases at Bode but not at Chanban. The growth of the
784 boundary layer after sunrise and entrainment of air from the free troposphere, with lower CO
785 mixing ratios, causes CO to decrease sharply during the day at Bode. At Chanban, on the other
786 hand, since the mixing ratios are already more representative of the local and regional
787 background levels which will also be prevalent in the lower free troposphere, CO does not
788 decrease notably during the daytime growth of the boundary layer as observed at Bode.

789 Similarly, while there is very little diurnal variation in the CH_4 mixing ratios at Chanban, there is
790 a strong diurnal cycle of CH_4 at Bode, similar to CO_2 there. At Chanban, the CH_4 mixing ratio
791 only shows a weak minimum at around 11 am, a slow increase during the day until a its peak
792 around 22:00, followed by a slow decrease during the night and a more rapid decrease through
793 the morning. The cause of this diurnal pattern at Chanban is presently unclear, but ~~it is clear that~~
794 the levels could be ~~are generally~~ representative of the regional background throughout the day
795 and show only limited influences of local emissions.

796   4. **Conclusions**

797   A cavity ring down spectrometer (Picarro G2401, USA) was used to measure ambient CO_2, CH_4,

798   CO, and water vapor mixing ratios at a semi-urban site (Bode) in the Kathmandu Valley for a

799   year. This was the first 12-months of continuous measurements of these four species in the

800   Kathmandu Valley in the foothills of the central Himalaya. Simultaneous measurement was

801   carried out at a rural site (Chanban) for approximately 3 months to evaluate urban-rural

802   differences.

803   The measurement also provided an opportunity to establish diurnal and seasonal variation of

804   these species in one of the biggest metropolitan cities in the foothills of Himalayas. Annual

805   average of the mixing ratio of CH_4 and CO_2 in Bode revealed that they were higher than the

806   ~~concentrations~~ mixing ratios at the background sites such as the Mauna Loa, USA and Mt.

807   Waliguan, China, as well as higher than urban/semi-urban sites in nearby regions such as

808   Ahmedabad and Shadnagar in India~~, and Nanjing in China~~. These comparisons highlight

809   potential sources of CH_4 and CO_2 in the Kathmandu Valley, such as brick kilns in the valley.

810   Polluted air masses were transported to the site mainly by two major local wind circulation

811   patterns, East-South/North East and West-Southwest throughout the observation period. Strong

812   seasonality was observed with CO compared to CO_2 and CH_4. Winter and pre-monsoon high CO

813   are linked to emission sources active in these seasons only and are from east-southeast and west-

814   southwest. Emission from the east-southeast are most likely related to brick kilns (winter and

815   pre-monsoon), which are in close proximity to Bode. Major city-centers are located in the west-

816   southwest of Bode (vehicular emission) which impact the site all-round the year, although higher

817   during winter season. Winter high was also observed with CO_2 and CH_4, which are mostly local

818   influence of brick kilns, trash burning and emission from city-center. Nighttime and early

819   morning accumulation of pollutants in winter due to a shallow stable mixing height (ca. 200 m)

820   also contribute to elevated levels than other seasons. ~~Regional transport into the Kathmandu~~

821   ~~Valley could be related to CO_2 peak during pre-monsoon. The highest CH_4 during the post-~~

822   ~~monsoon could be associated with agricultural activity northeast of Bode.~~ Diurnal variation

823   across all seasons indicates the influence of rush-hour emissions related to vehicles and

824 residential emissions. The evolution of the mixing layer height (200-1200 m) was a major factor

825 which controls the morning-evening peak, afternoon low and night-early morning accumulation

826 or decay. Thus the geographical setting of the Kathmandu Valley and its associated meteorology

827 play a key role in the dispersion and ventilation of pollutants in the Kathmandu Valley. The ratio

828 of CO/CO_2 across different season and wind direction showed that emissions from inefficient

829 gasoline vehicles, brick kilns, residential cooking and diesel combustion are likely to impact

830 Bode.

831 The differences in mean values for urban-rural measurements at Bode and Chanban is highest for

832 CO (64 %) compared to CO_2 (3.8%) and CH_4 (12%). Low values of CH_4 and CO_2 mixing ratios

833 at the Chanban site could represent ~~a~~ regional background mixing ratios.

834 This study has provided valuable information on key greenhouse gases and air pollutants in the

835 Kathmandu Valley and the surrounding regions. These observations can be useful as ground-

836 truthing for evaluation of satellite measurements, as well as climate and regional air quality

837 models. The overall analysis presented in the paper will contribute along with other recent

838 measurement and analysis to ~~can~~ providing ~~e~~ a sound scientific basis for reducing emissions of

839 greenhouse gases and air pollutants in the Kathmandu Valley.

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Table 1. Instruments and sampling at Bode (semi-urban site) and Chanban (rural site)

Site	Instrument	Species	sampling interval	Measurement period	inlet/sensor height above ground (m)
Bode	i. Cavity ring down spectrometer (Picarro G2401, USA)	CO_2, CH_4, CO, water vapor	5 sec	06 Mar 2013 - 05 Mar 2014 14 Jul 2015 - 07 Aug 2015	20
	ii. CO monitor (Horriba AP370, USA)	CO	5 min	06 Mar 2013 – 07 June 2013	20
	iii. Ceilometer (Vaisala CL31, Finland		15-52 min	06 Mar 2013 – 05 Mar 2014	15
	iv. AWS (Campbell Scientific, USA)		1 min		23
	a. CS215	RH, T		06 Mar 2013 – 24 Apr 2013	
	b. CS300 Pyranometer	SR		06 Mar 2013 - 05 Mar 2014 14 Jul 2015 - 07 Aug 2015	
	c. RM Young 05103-5	WD, WS		06 Mar 2013 - 05 Mar 2014 14 July 2015 - 07 Aug 2015	
	v. Airport AWS (Environdata, Australia)				
	a. TA10	T		18 Jun 2013 – 13 Jan 2013	
	b. RG series	RF		06 Mar 2013 – 15 Dec 2013	
Chanban	i. Cavity ring down spectrometer (Picarro G2401, USA)	CO_2, CH_4, CO, water vapor	5 sec	15 July 2015 - 03 Oct 2015	3
	ii AWS (Davis Vantage Pro2, USA)	RH, T, SR, WD, WS, RF, P	10 min	14 July 2015 - 07 Aug 2015	2

AWS: Automatic weather station, RH: ambient relative humidity, T: ambient temperature, SR: global solar radiation, WS: wind speed, WD: wind direction, RF: rainfall, P: ambient pressure

Table 2. Summary of monthly average CH_4 and CO_2 mixing ratios observed at Bode, a semi-urban site in the Kathmandu Valley during March 2013 to Feb 2014 [mean, standard deviation (SD), median, minimum (Min.), maximum (Max.) and number of data points of hourly average values]

Month	CH₄ (ppm)					CO₂ (ppm)					Data points
	Mean	SD	Median	Min.	Max.	Mean	SD	Median	Min.	Max.	
Mar	2.207	0.245	2.152	1.851	3.094	426.6	26.4	418.3	378.8	510.8	596
Apr	2.183	0.252	2.094	1.848	3.121	430.3	27.4	421.0	397.0	536.9	713
May	2.093	0.174	2.040	1.863	2.788	421.7	22.1	413.4	395.9	511.2	725
Jun	2.061	0.142	2.017	1.869	2.675	417.9	21.3	410.4	390.5	495.7	711
Jul	2.129	0.168	2.074	1.893	2.770	410.3	18.2	406.3	381.0	471.0	500
Aug	2.274	0.260	2.181	1.953	3.219	409.9	22.8	405.3	376.1	493.1	737
Sep	2.301	0.261	2.242	1.941	3.331	414.9	30.2	404.0	375.9	506.2	710
Oct	2.210	0.195	2.156	1.927	2.762	417.0	25.1	411.8	381.9	486.7	743
Nov	2.207	0.203	2.178	1.879	2.705	417.2	20.7	415.7	385.7	478.9	717
Dec	2.206	0.184	2.193	1.891	2.788	417.7	17.3	418.0	386.7	467.6	744
Jan	2.233	0.219	2.198	1.889	2.744	424.8	20.9	422.3	392.7	494.5	696
Feb	2.199	0.223	2.152	1.877	2.895	423.2	22.0	417.9	392.2	484.6	658
Annual	2.19~~23~~	0.066~~22~~ 4	2.140~~34~~	1.848	3.331	419.3 ~~9~~	~~6~~23.0	41~~3~~4.7 ~~0~~	375.9	536.9	~~83~~53

Table 3. Summary of CH_4 and CO_2 mixing ratios at Bode across four seasons during March 2013 to Feb 2014 [seasonal mean, one standard deviation (SD), median, minimum (Min.) and maximum (Max.)]

Season	CH_4 (ppm)						CO_2 (ppm)				
	Mean	SD	Median	Min.	Max.		Mean	SD	Median	Min.	Max.
Pre-Monsoon	2.157	0.230	2.082	1.848	3.121		426.2	25.5	417.0	378.8	536.9
Monsoon	2.199	0.241	2.126	1.869	3.331		413.5	24.2	407.1	375.9	506.2
Post-Monsoon	2.210	0.200	2.167	1.879	2.762		417.3	23.1	414.1	381.9	486.7
Winter	2.214	0.209	2.177	1.877	2.895		421.9	20.3	419.3	386.7	494.5

Table 4. Comparison of monthly average CH_4 and CO_2 mixing ratios at a semi-urban and a rural site in Nepal (this study) with other urban and background sites in the region and elsewhere.

Site	Bode, Nepal				Chanban, Nepal		Mauna Loa, USA		Waliguan, China	
Setting	(Urban)				(Rural)		(Background)[c]		(Background)[d]	
Species	CO_2	CH_4	$*CO_2$	$*CH_4$	$*CO_2$	$*CH_4$	CO_2	CH_4	CO_2	CH_4
Unit	ppm	ppm	ppm	ppm	ppm	ppm	ppm	Ppm	ppm	ppm
Mar 2013	426.6	2.207					397.3	1.840	399.5	1.868
Apr	430.3	2.183					398.4	1.837	402.8	1.874
May	421.7	2.093					399.8	1.834	402.5	1.878
Jun	417.9	2.061					398.6	1.818	397.4	1.887
Jul	410.3	2.129					397.2	1.808	393.3	1.888
Aug	409.9	2.274	411.3	2.281	403.4	2.050	395.2	1.819	392.0	1.893
Sep	414.9	2.301	419.9	2.371	399.1	2.102	393.5	1.836	393.1	1.894
Oct	417.0	2.210					393.7	1.836	395.6	1.876
Nov	417.2	2.207					395.1	1.835	397.1	1.875
Dec	417.7	2.206					396.8	1.845	398.6	1.880
Jan 2014	424.8	2.234					397.8	1.842	398.8	1.865

Feb	423.2	2.199		397.9	1.834	401.1	1.878
Annual							
Bode	419.3	2.192					
Mauna Loa				396.8	1.832		
Waliguan						397.7	1.880
~~Nanjing (2011)[a]~~	~~406.5~~						
Shadnagar (2014)[a]	394.0						
Ahemadabad (2013-2015)[b]	413.0	1.920					

*The monthly values for CO_2 and CH_4 in 2015, [a] ~~Huang et al., 2015,~~ [a]Sreenivas et al., 2016, [b]Chandra et al., 2016, [c]Dlugokencky et al., 2017; NOAA, 2015, [d]Dlugokencky et al., 2016; Dlugokencky et al., 2016a.

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[d]~~ftp://aftp.cmdl.noaa.gov/data/trace_gases/co2/flask/surface/wlg/; ftp://aftp.cmdl.noaa.gov/data/trace_gases/ch4/flask/surface/wlg/~~

Table 5. Emission ratio of CO/CO_2 (ppb ppm^{-1}) derived from emission factors (gram of gas emitted from per kilogram of fuel burned, except transport sector which is derived from gram of gases emitted per kilometer distance travelled)

Sectors	Details	CO/CO_2	Reference
1. Residential/Commercial			
i. LPG		4.8	Smith et al. (2000)
ii. Kerosene		13.4	Smith et al. (2000)
iii. Biomass		52.9 - 98.5	*
iv. Diesel power			The World Bank
generators	< 15 year old	5.8	(2014)
	>15 year old	4.5	
2. Transport			**
a. Diesel			
i. HCV diesel bus	>6000cc, 1996-2000	4.9	
	post 2000 and 2005	5.4	
ii. HCV diesel truck	>6000cc, post 2000	7.9	
b. Petrol			
i. 4 stroke motorcycle	<100 cc, 1996-2000	68	
	100-200 cc, Post 2000	59.6	
ii. Passenger cars	<1000 cc, 1996-2000	42.4	
iii. Passenger cars	<1000 cc, Post 2000	10.3	
3. Brick industries			
i. BTK fixed kiln		17.2	Weyant et al. (2014)
ii. Clamp brick kiln		33.7	Stockwell et al. (2016)
iii. Zigzag brick kiln		3.9	Stockwell et al. (2016)
4. Open burning			
i. Mixed garbage		46.9	Stockwell et al. (2016)
ii. Crop-residue		51.6	Stockwell et al. (2016)

* Westerdahl et al. (2009)

** http://www.cpcb.nic.in/Emission_Factors_Vehicles.pdf

Table 6. ~~Seasonal a~~Average (SD) of the ratio of dCO to dCO_2, their Geometric mean (GeoSD) over a period of 3 hours during (a) morning peak ~~and~~ (b) evening peak and (c) seasonal (all hours) of~~in~~ the ambient mixing ratios of CO and CO_2. And their lower and upper bound (LB and UB).

Period	Season	Mean (SD)	Median	N	Geomean (GeoSD)	LB	UB
a. Morning hours (7:00-9:00)	Pre-monsoon	7.6 (3.1)	7.8	249	11.3 (1.5)	5.2	24.8
	Monsoon	2.2 (1.6)	1.9	324	9.9 (1.9)	2.7	36.3
	Post-monsoon	3.1 (1.4)	2.8	183	11.1 (1.5)	4.7	26.3
	Winter*	11.2 (4.4)	11	255	11.4 (1.5)	5.3	24.2
b. Evening hours (19:00-21:00)	Pre-monsoon	15.1 (9.0)	12.7	248	10.5 (1.7)	3.5	31.6
	Monsoon	8.0 (5.2)	6.3	323	10.2 (1.8)	3.1	33.5
	Post-monsoon	11.5 (5.6)	10.6	182	11.0 (1.6)	4.4	27.6
	Winter	21.6 (14.1)	18.2	254	10.2 (1.8)	3.1	33.6
c. Seasonal (all hours)	Pre-monsoon	12.2 (13.3)	8.8	1740	8.2 (2.4)	1.4	48.4
	Monsoon	7.5 (13.5)	2.9	2176	5.9 (3.3)	0.5	65.6
	Post-monsoon	8.3 (12.4)	4.4	1289	6.8 (3.0)	0.8	59.2
	Winter	15.1 (13.3)	12.5	1932	9.2 (2.1)	2.0	41.7

*The morning peak was one hour delayed in winter, thus the 8:00-10:00 period data was used in the analysis.
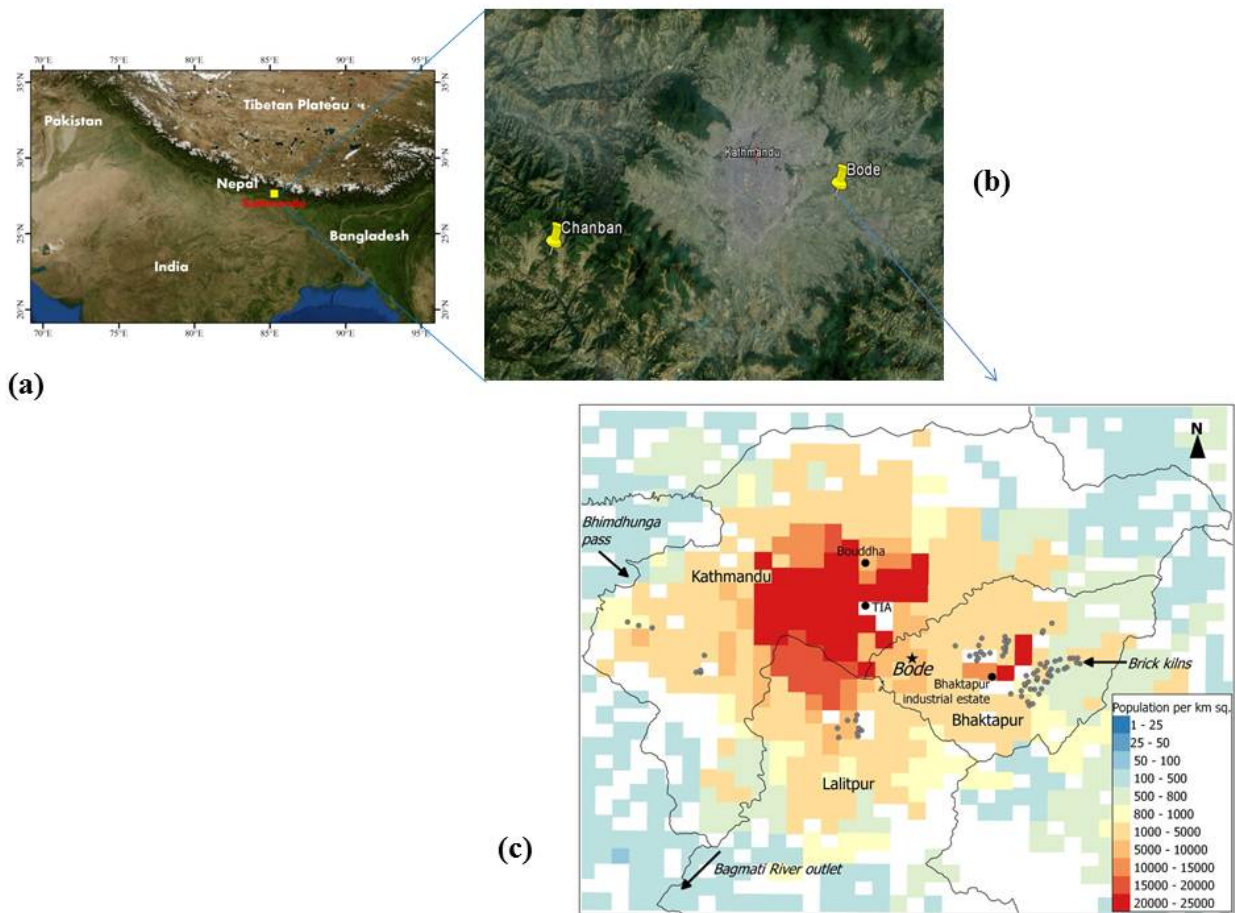
Figure 1. Location of measurement sites: (a) Kathmandu Valley (b) semi-urban measurement site at Bode in Kathmandu Valley, and a rural measurement site at Chanban in Makawanpur district Nepal, (c) general setting of Bode site. Colored grid and TIA represent population density and the Tribhuvan International Airport, respectively.
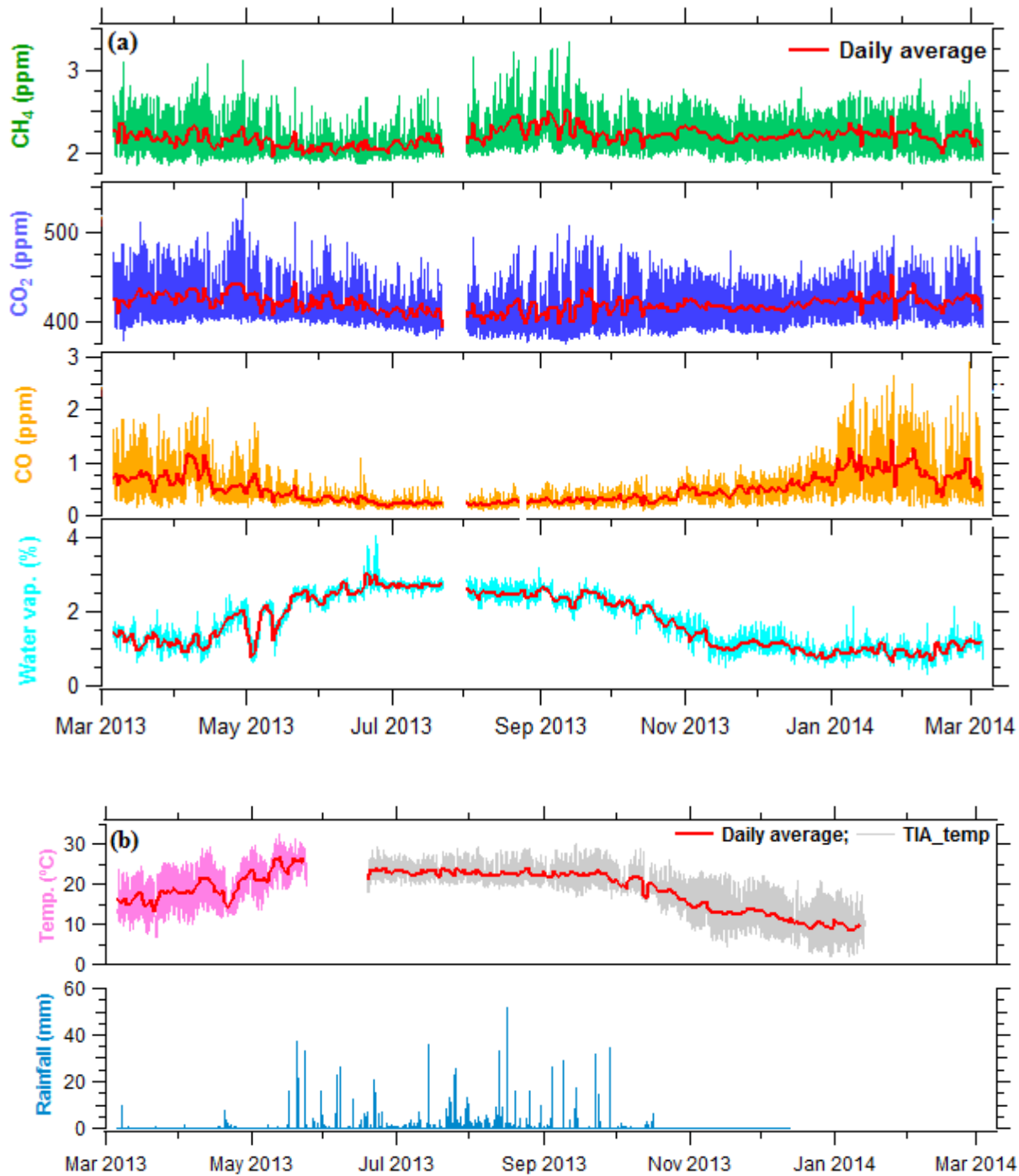
Figure 2. Time series of hourly average (a) mixing ratios of CH_4, CO_2, CO, and water vapor measured with a cavity ring down spectrometer (Picarro G2401) at Bode, and (b) temperature and rainfall monitored at the Tribhuvan International Airport (TIA), ~4 km to the west of Bode site in the Kathmandu Valley, Nepal. Temperature shown in pink color is observed at Bode site.
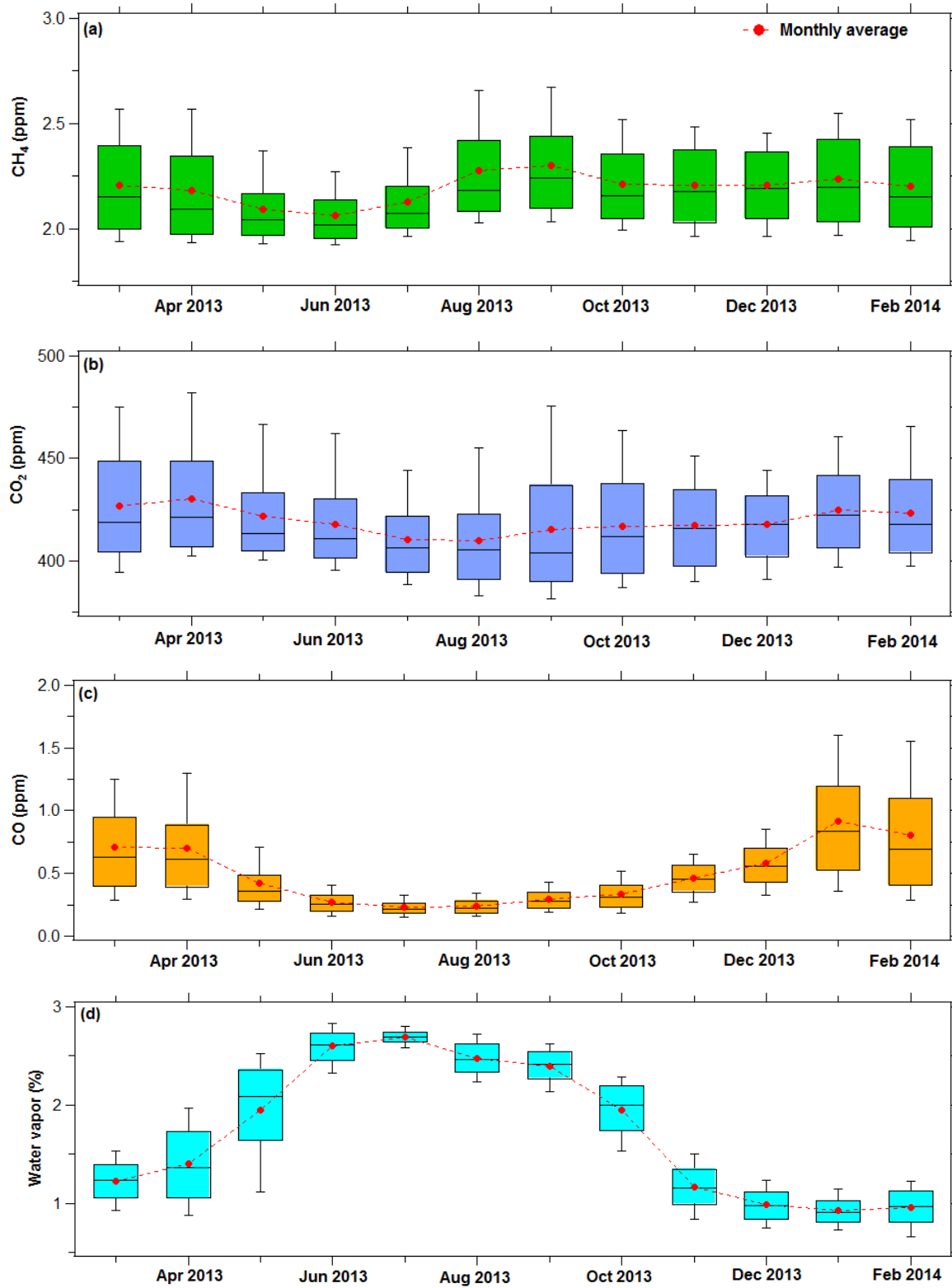
Figure 3. Monthly variations of the mixing ratios of hourly (a) CH_4, (b) CO_2, (c) CO, and (d) water vapor observed at a semi-urban site (Bode) in the Kathmandu Valley over a period of a year. The lower end and upper end of the whisker represents 10th and 90th percentile, respectively; the lower end and upper end of each box represents 25th and 75th percentile, respectively, and black horizontal line in the middle of each box is the median for each month while red dot represents mean for each month.

(a)
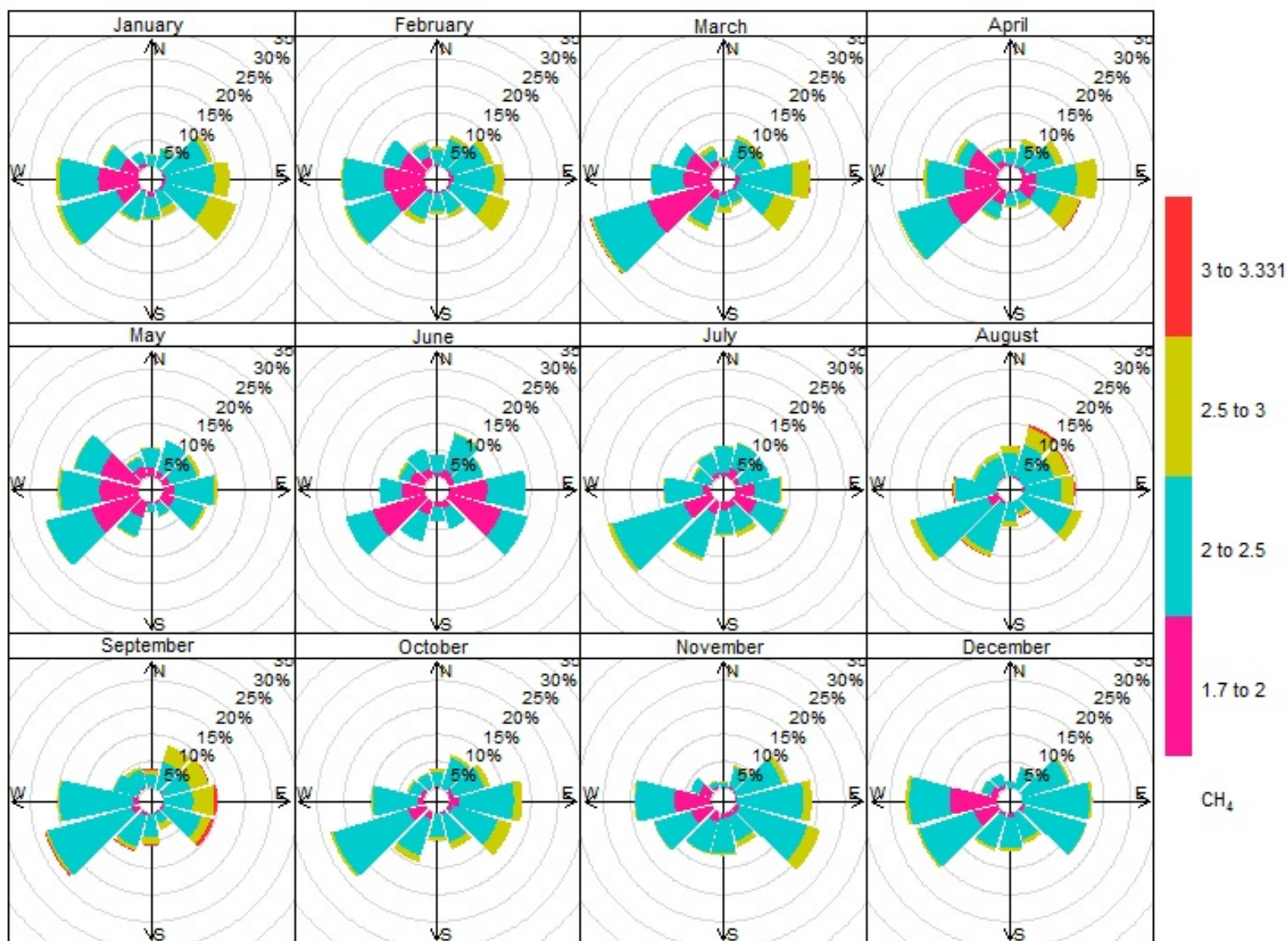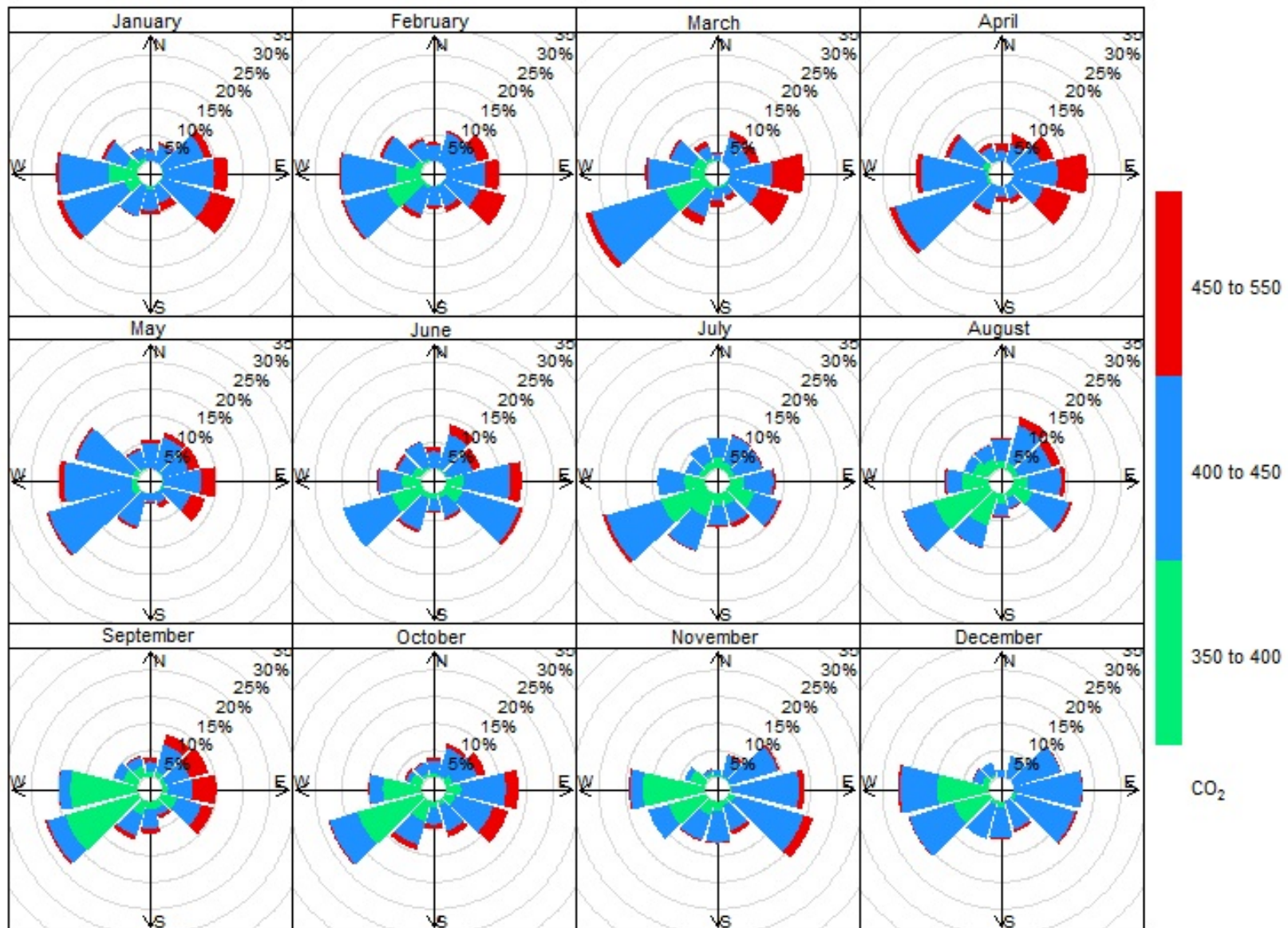
(c)
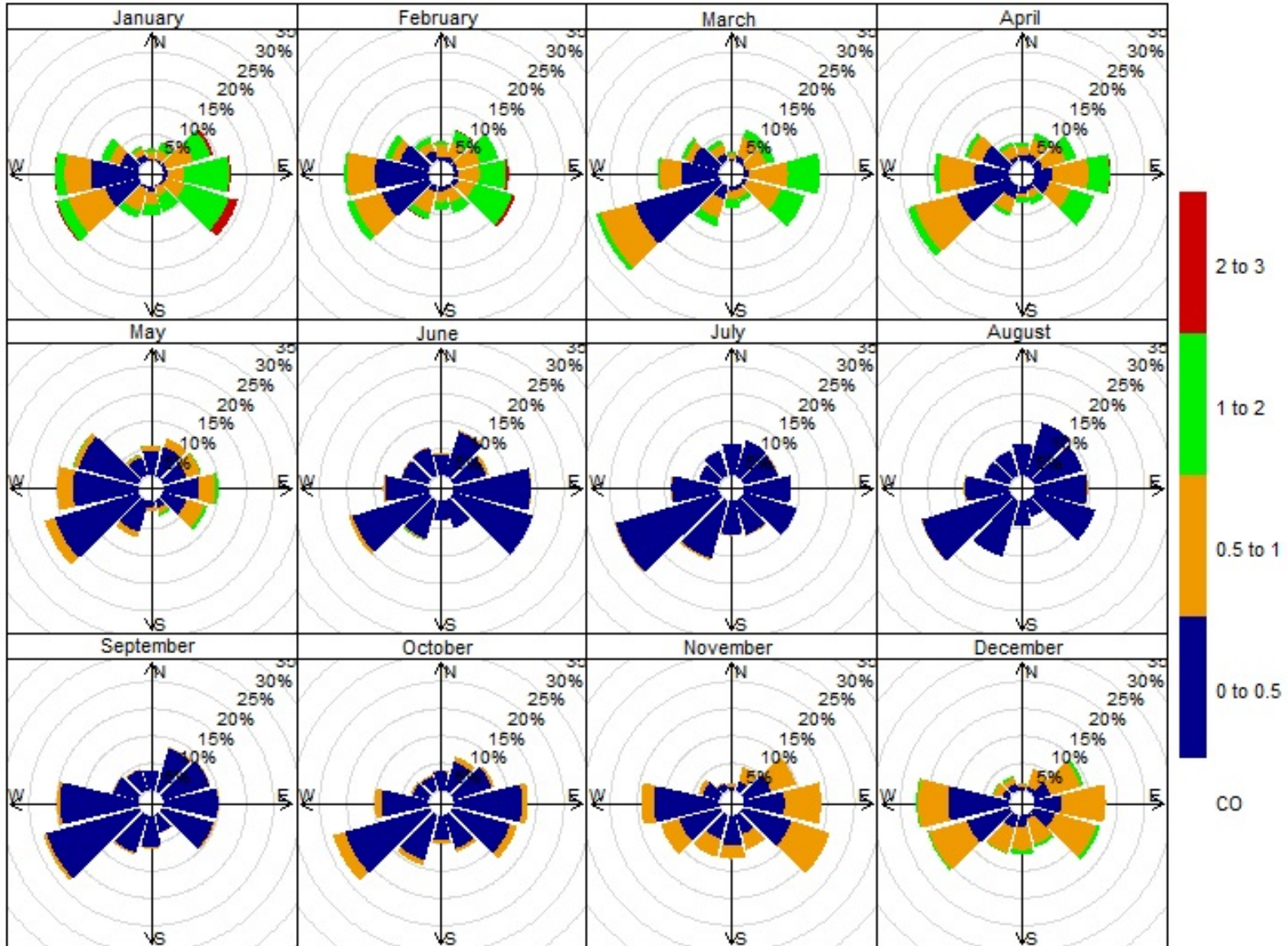
Figure 4. Relation between ~~Pollution rose of the hourly CH₄ and CO₂~~ mixing ratios and wind direction observed at Bode in the Kathmandu Valley (a) CH₄, ~~and~~ (b) CO₂ and (c) CO from March 2013 to February 2014. ~~Pollution rose~~The figure shows variations of ~~pollutants~~ CH₄, CO₂ and CO mixing ratios based on frequency ~~of~~ counts of~~by~~ wind direction (in %) as represented by circle. The color represents the different mixing ratios of the gaseous species. The units of CH₄, ~~and~~ CO₂ and CO are in ppm.
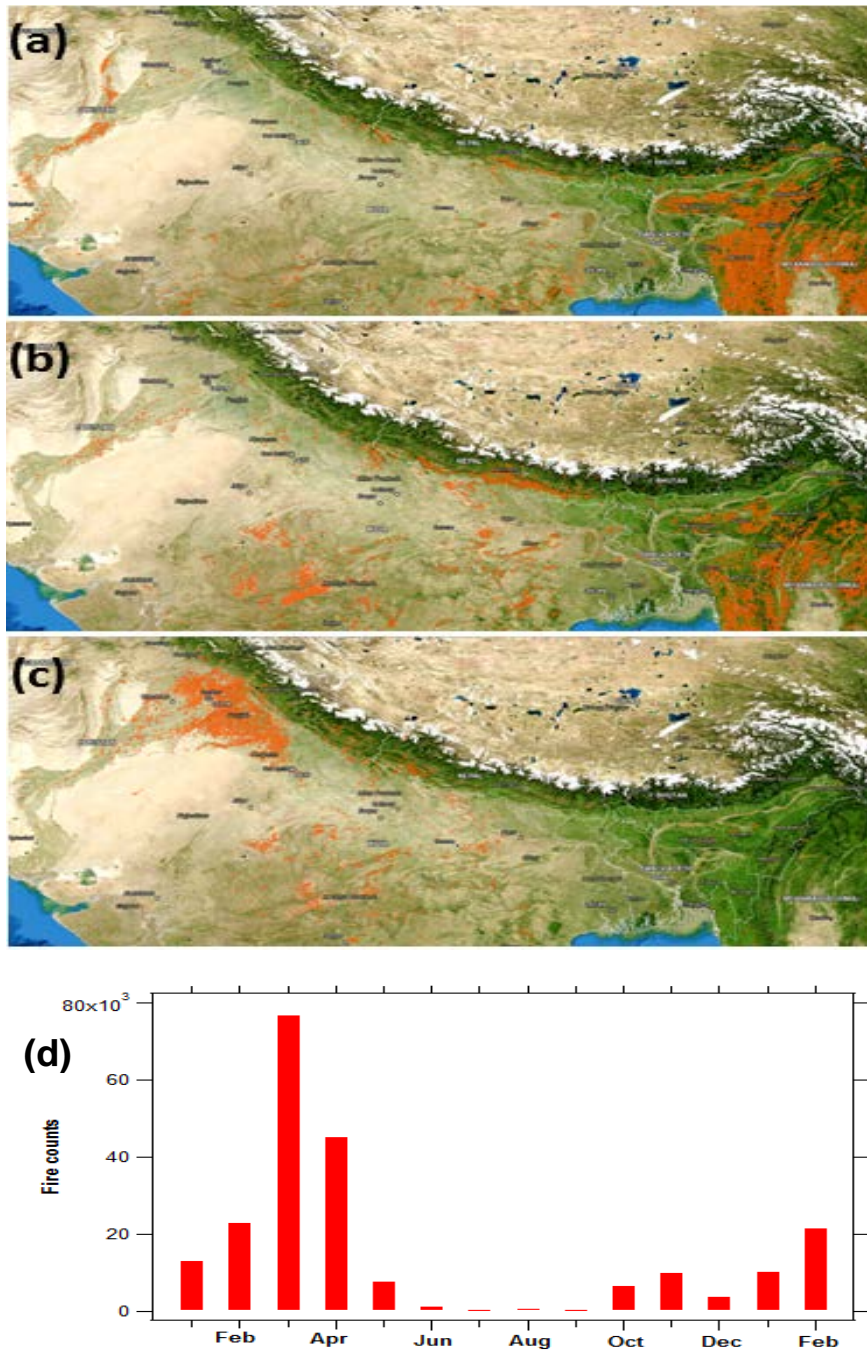
Figure 5. Satellite detected fire counts in (a) Mar, (b) Apr, (c) May 2013 in the broader region surrounding Nepal and (d) total number of fire counts detected by MODIS instrument onboard the Aqua satellite during Jan 2013-Feb 2014. Source: https://firms.modaps.eosdis.nasa.gov/firemap/
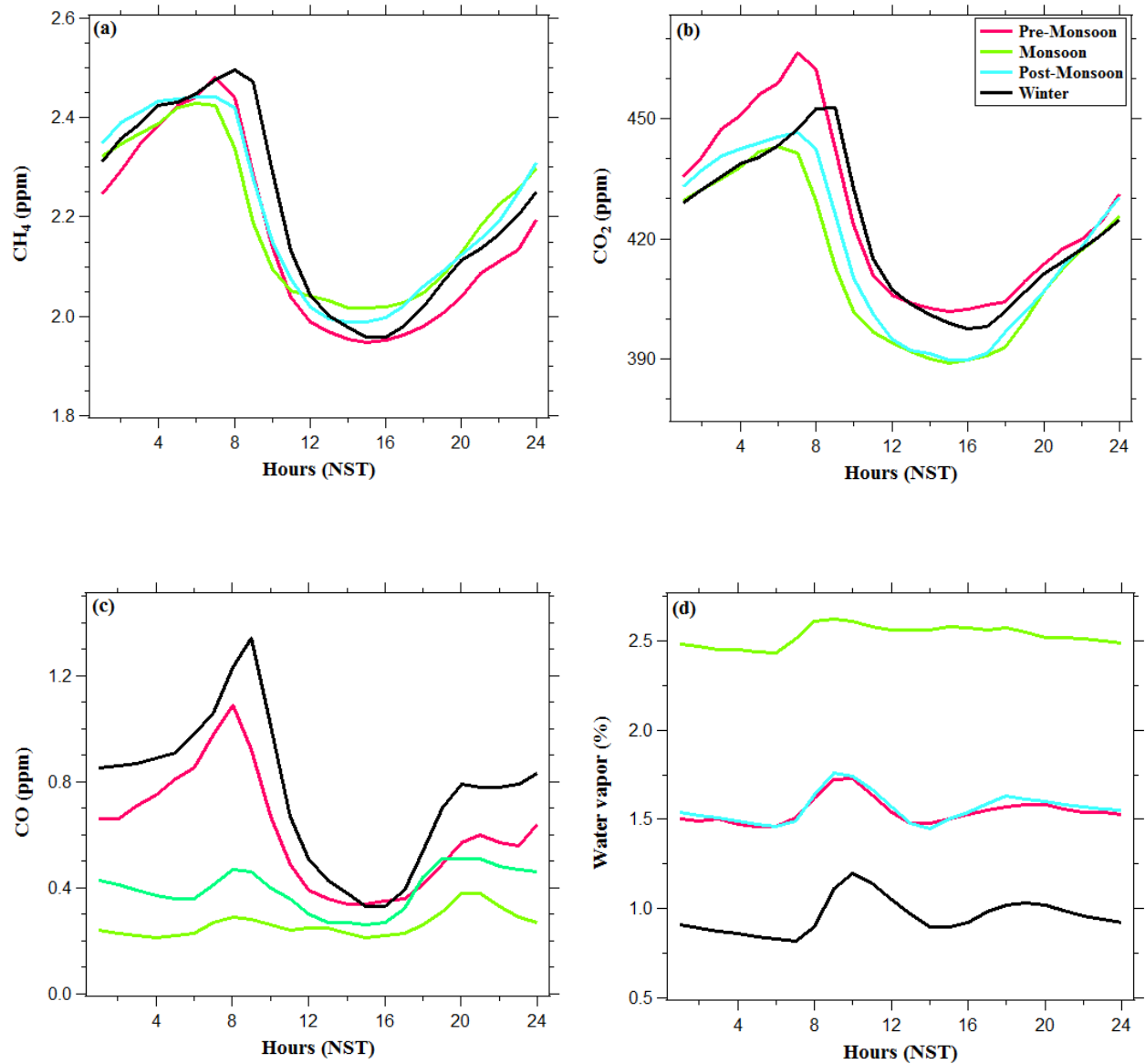
Figure 6. Diurnal variations of hourly mixing ratios in different seasons (a) CH_4, (b) CO_2, (c) CO, and (d) water vapor observed at Bode (semi-urban site) in the Kathmandu Valley during March 2013-February 2014. Seasons are defined as Pre-monsoon: Mar-May, Monsoon: Jun-Sep, Post-monsoon: Oct-Nov, Winter: Dec-Feb. The x axis is in Nepal Standard Time (NST).

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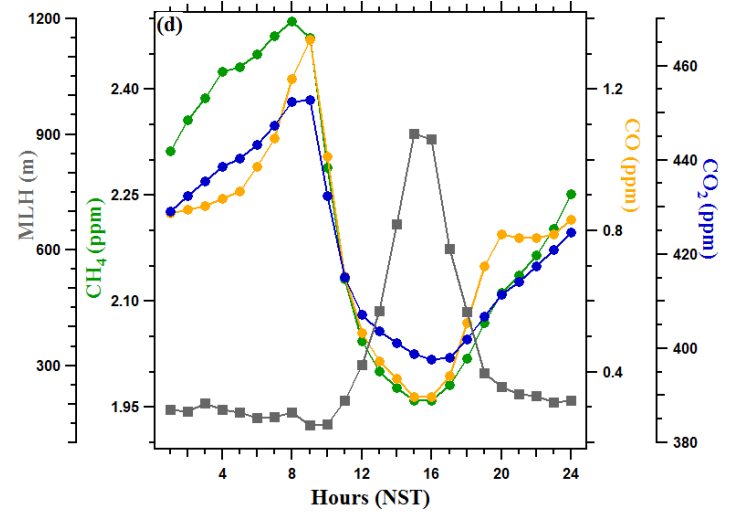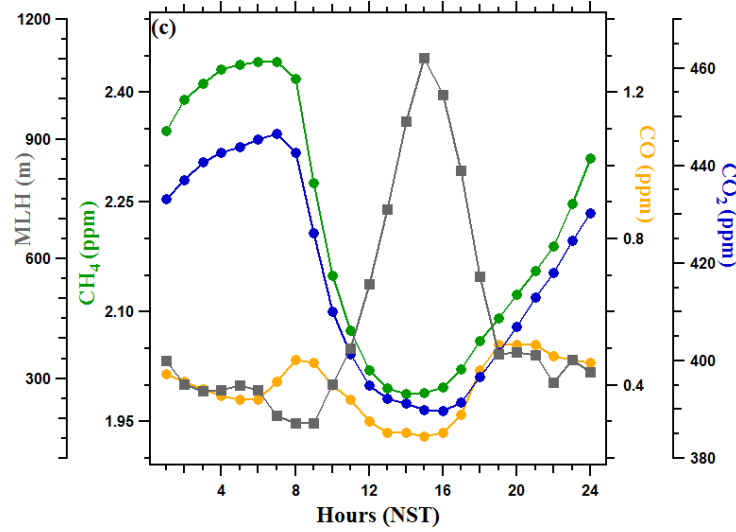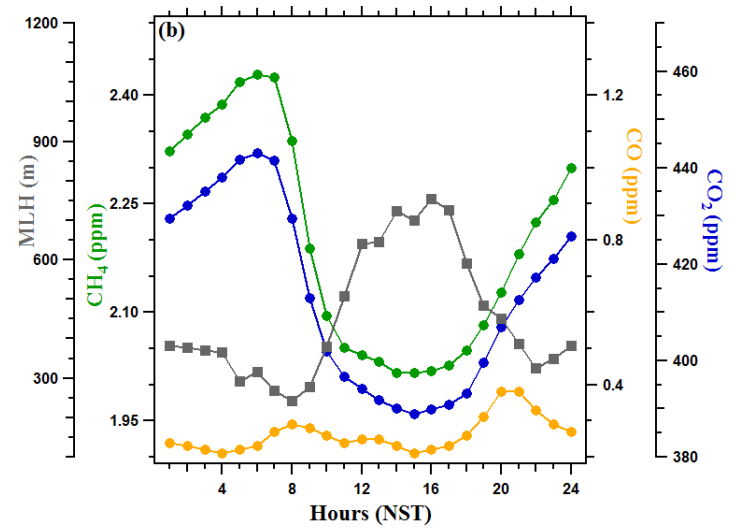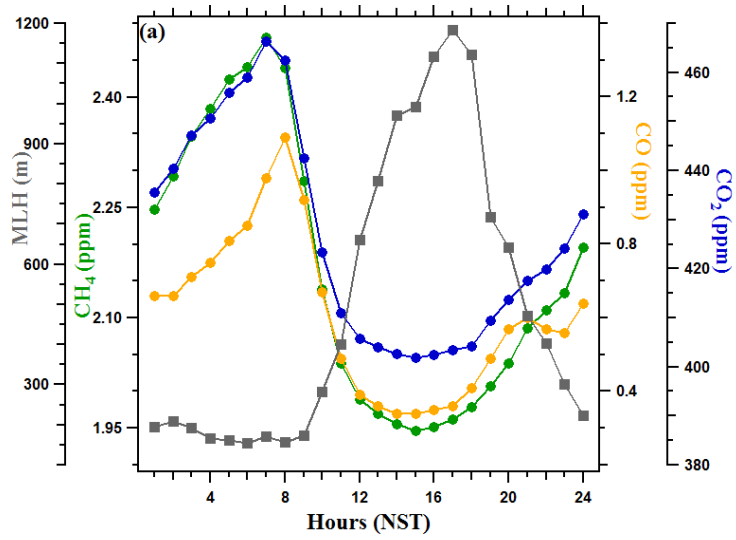
Figure 7. Diurnal variations of hourly mixing ratios of CH_4, CO_2, CO, and mixing layer height (MLH) at Bode (a semi-urban site in the Kathmandu Valley) in different seasons (a) pre-monsoon (Mar-May), (b) monsoon (Jun-Sep), (c) post-monsoon (Oct-Nov) and (d) winter (Dec-Feb) during March 2013- Feb 2014.
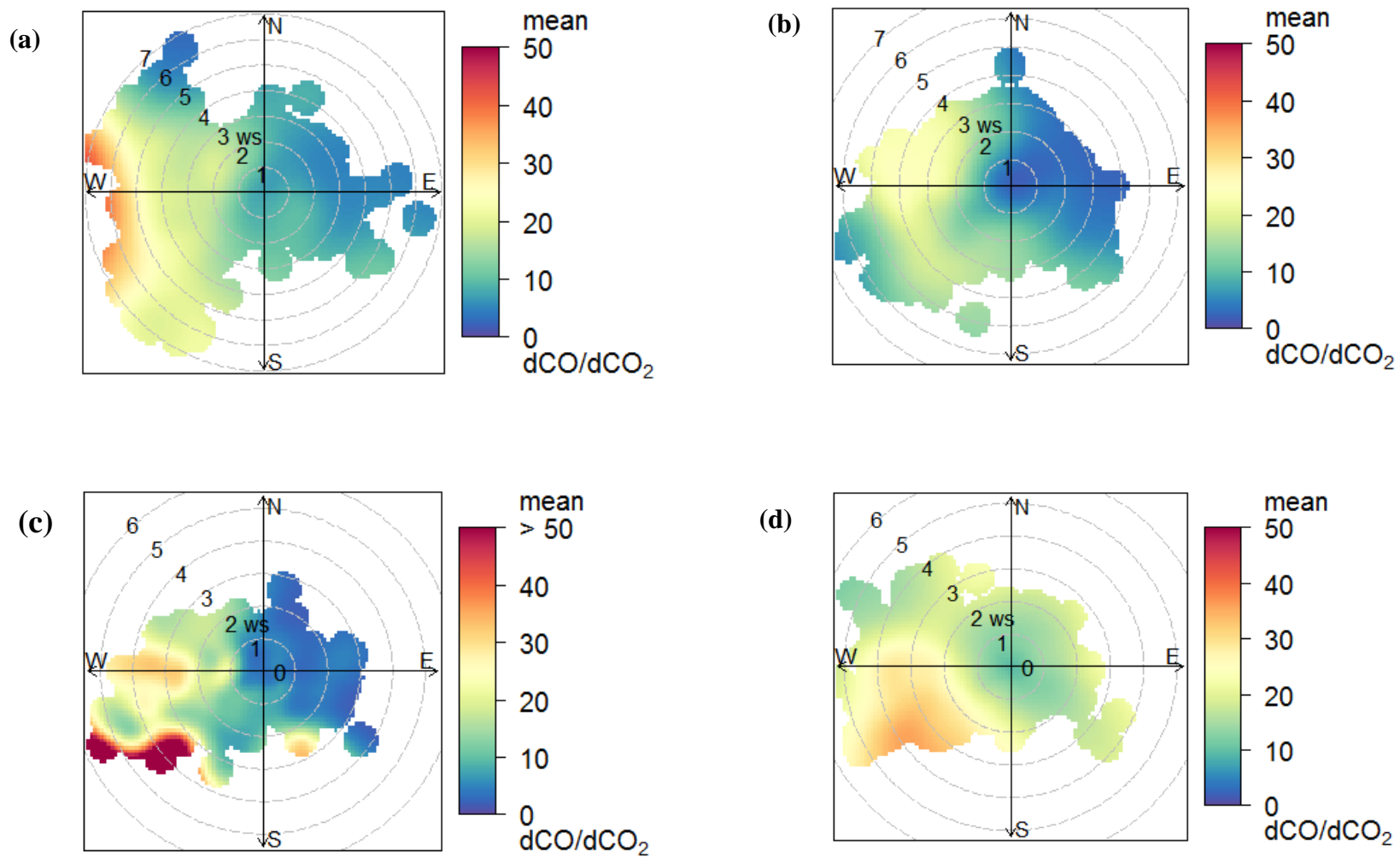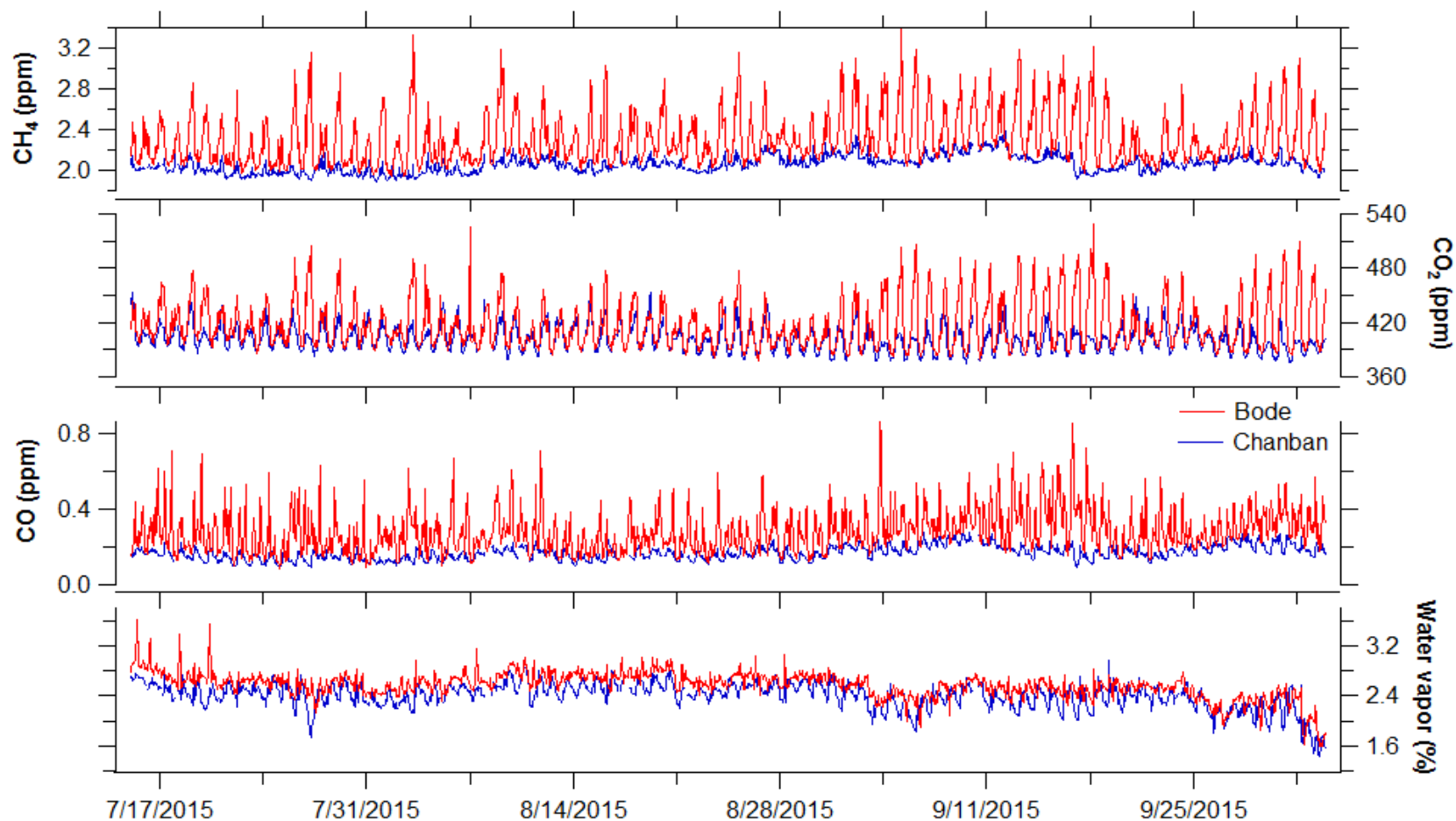
Figure 8. Seasonal polar plot of hourly dCO/dCO$_2$ ratio based upon wind direction and wind speed: (a) pre-monsoon, (b) monsoon, (c) post-monsoon and (d) winter seasons.
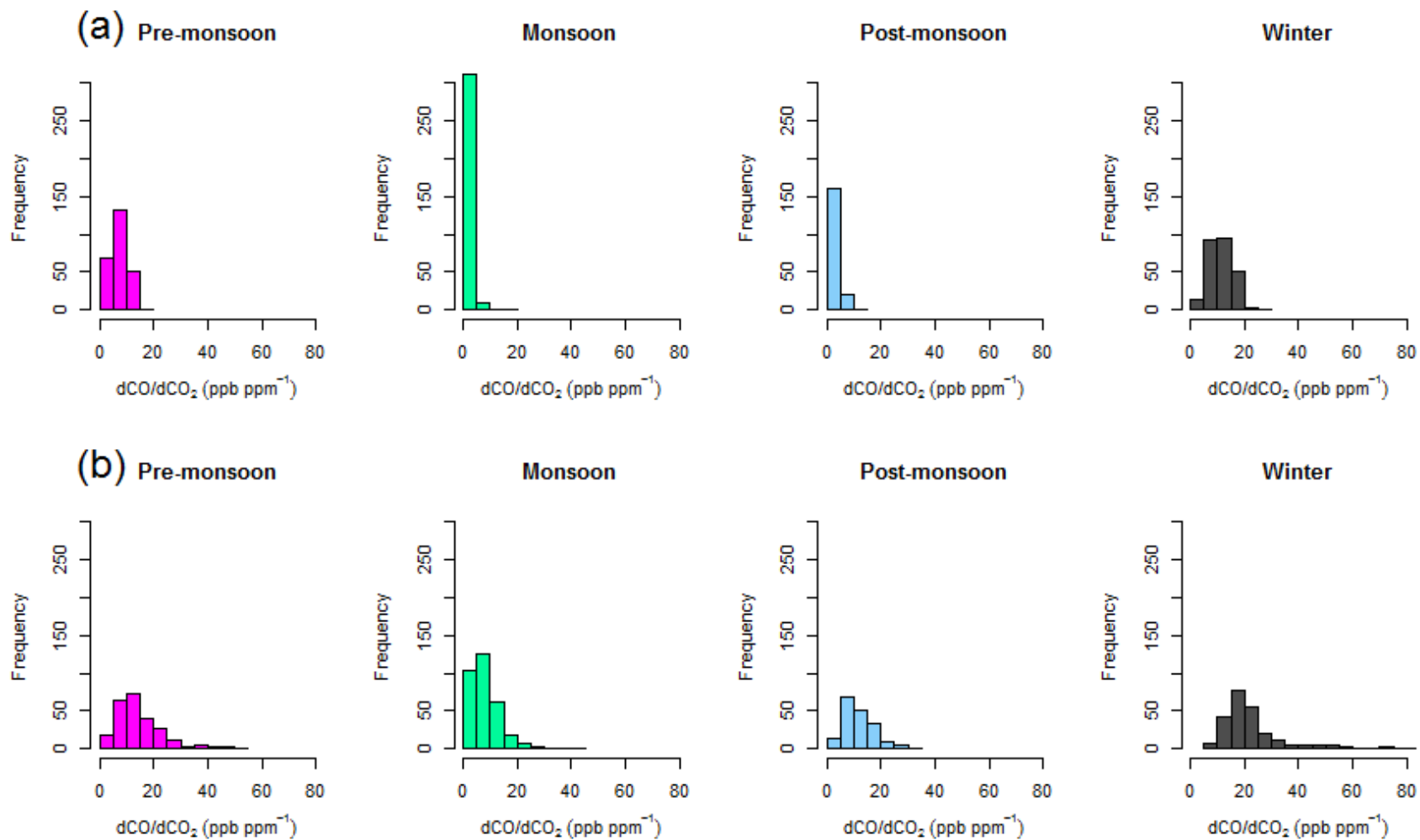
Figure 9. Seasonal frequency distribution of hourly dCO/dCO$_2$ ratio (a) morning hours (7:00-9:00) in all season except winter (8:00-10:00), (b) evening hours (19:00-21:00)

Figure 10. Comparison of hourly average mixing rations of CH_4, CO_2, CO, and water vapor observed at Bode (a semi-urban site) in the Kathmandu Valley and at Chanban (a rural/background site) in Makawanpur district, ~ 20 km from Kathmandu, on other side of a tall ridge.
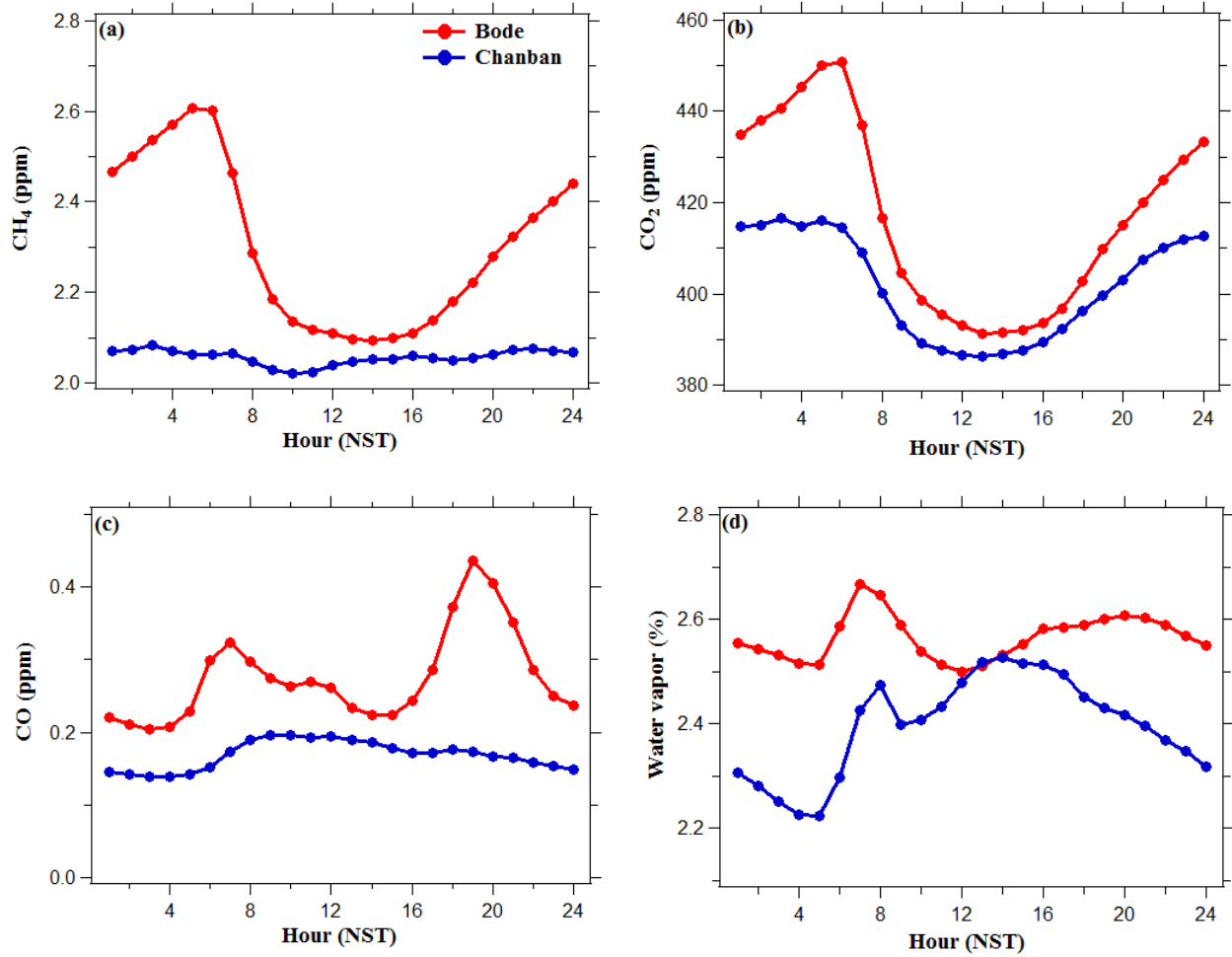
Figure 11. Diurnal variations of hourly average mixing ratios of (a) CH_4, (b) CO_2, (c) CO and (d) water vapor observed at Bode in the Kathmandu Valley and at Chanban in Makawanpur district during 15 July- 03 October 2015.