

1 Mixing state of oxalic acid containing particles in the rural area of Pearl
2 River Delta, China: implications for the formation mechanism of oxalic
3 acid

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42 **Abstract:**

43 The formation of oxalic acid and its mixing state in atmospheric particulate
44 matter (PM) were studied using a single particle aerosol mass spectrometer (SPAMS)
45 in the summer and winter of 2014 in Heshan, a supersite in the rural area of the Pearl
46 River Delta (PRD) region in China. Oxalic acid-containing particles accounted for 2.5%
47 and 2.7% in total detected ambient particles in summer and winter, respectively.
48 Oxalic acid was measured in particles classified as elemental carbon (EC), organic
49 carbon (OC), elemental and organic carbon (ECOC), biomass burning (BB), heavy
50 metal (HM), secondary (Sec), sodium-potassium (NaK) and dust. Oxalic acid was
51 found predominantly mixing with sulfate and nitrate during the whole sampling
52 period, likely due to aqueous phase reactions. In summer, oxalic acid-containing
53 particle number and ozone concentration followed a very similar trend, which may
54 reflect the significant contribution of photochemical reactions to oxalic acid formation.
55 The HM type particles were the most abundant oxalic acid particles in summer and
56 the diurnal variations of peak area of iron and oxalic acid show opposite trends, which
57 suggest a possible loss of oxalic acid through the photolysis of iron oxalato complexes
58 during the strong photochemical activity period. In wintertime, carbonaceous type
59 particles contained a substantial amount of oxalic acid as well as abundant carbon
60 clusters and biomass burning markers. The general existence of nitric acid in oxalic
61 acid-containing particles indicates an acidic environment during the formation process
62 of oxalic acid. The peak areas of nitrate, sulfate and oxalic had similar temporal
63 change in the carbonaceous type oxalic acid particles, and the
64 organosulfate-containing oxalic acid particles well correlated with total oxalic acid
65 particles during the episode, which suggests the formation of oxalic acid is closely
66 associated with the oxidation of organic precursors in aqueous phase.

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68 **Keywords:** Oxalic acid; Single particles; Mixing state; Photochemical process;
69 Aqueous phase reactions.

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71 **1. Introduction**

72 Organic aerosol, typically a large fraction of fine particles, contains more than
73 thousands of organic compounds and contributes to visibility reduction,
74 photochemical smog, climate change and adverse health effects (Novakov and Penner,
75 1993;Goldstein and Galbally, 2007;Jimenez et al., 2009;Poschl and Shiraiwa, 2015). A
76 significant component of organic aerosol is secondary organic aerosol (SOA) formed
77 from the gas phase oxidation of volatile organic compounds (VOCs) followed by
78 partitioning of products into particles or from heterogeneous reactions of VOCs with
79 particles (Hallquist et al., 2009;Zhang et al., 2015). Dicarboxylic acids (DCAs) are
80 abundant and ubiquitous constituents in SOA and can be effective tracers for the
81 oxidative processes leading to the formation of SOA (Kawamura and Ikushima,
82 1993;Ervens et al., 2011;Wang et al., 2012;Cheng et al., 2013). DCAs normally have
83 high water solubility and low vapor pressure, so they play important roles in
84 controlling the hygroscopic properties of organic aerosols (Prenni et al., 2003;Ma et
85 al., 2013) and activating cloud condensation nuclei (Booth et al., 2009). The primary
86 emissions of DCAs from anthropogenic sources in urban areas are minor (Huang and
87 Yu, 2007;Stone et al., 2010), and they are mainly derived from secondary oxidation of
88 VOCs and subsequent intermediates (Ho et al., 2010;Myriokefalitakis et al., 2011).
89 High concentrations of DCAs have been observed in biomass burning plume (Kundu
90 et al., 2010;Kawamura et al., 2013) with more than 70% of DCAs produced from
91 photochemical oxidation of water-soluble organic compounds, and only a small
92 contribution from direct biomass burning emission (van Pinxteren et al., 2014).

93 The production of DCAs through photochemical reactions has been reported in
94 many field studies via the analysis of the diurnal and seasonal variations of
95 DCA(Kawamura and Ikushima, 1993;Kawamura and Yasui, 2005;Aggarwal and
96 Kawamura, 2008;Pavuluri et al., 2010;Ho et al., 2011;Wang et al., 2017), but the
97 mechanism of DCAs formation is still not well understood. Oxalic acid is usually the
98 most abundant DCA observed in the field (Kawamura et al., 2004;Ho et al.,
99 2007;Kawamura et al., 2010). A number of ground based and airborne field studies

100 have found a tight correlation between oxalic acid and sulfate in ambient particles and
101 cloud droplets, relating aqueous phase chemistry to the formation of oxalic acid in
102 aerosols and cloud droplets (Yao et al., 2002; Yao et al., 2003; Yu et al.,
103 2005; Sorooshian et al., 2006; Sorooshian et al., 2007a; Sorooshian et al.,
104 2007b; Miyazaki et al., 2009; Wonaschuetz et al., 2012; Wang et al., 2016). In recent
105 years, several model and laboratory studies suggested that the aqueous phase
106 oxidation of highly water-soluble organics like glyoxal, methylglyoxal and glyoxylic
107 acid can efficiently produce oxalic acid in aerosol particles and cloud droplets (Lim et
108 al., 2010; Myriokefalitakis et al., 2011; Ervens et al., 2014; Yu et al., 2014; McNeill,
109 2015). Recent stable carbon isotope studies and field observations have also suggested
110 that oxalic acid forms through aqueous phase reactions (Wang et al., 2012; Cheng et al.,
111 2015). However, the formation process of oxalic acid in ambient aerosols is still
112 associated with great uncertainty due to the oxidation rates of precursors and oxidant
113 levels in photochemistry and aqueous phase chemistry, which needs to be further
114 studied.

115 Online measurements of the size distribution of oxalic acid-containing particles
116 and the mixing state of oxalic acid with other compounds in aerosols are useful to
117 examine the formation and evolution of oxalic acid and SOA particles. Sullivan and
118 Prather investigated the diurnal cycle and mixing state of DCA-containing particles in
119 Asian aerosol outflow using aerosol time-of-flight mass spectrometry (ATOFMS), and
120 proposed the formation of DCA on Asian dust (Sullivan and Prather, 2007a). In
121 addition, Yang et al. (2009) measured oxalic acid particles in Shanghai and proposed
122 that in-cloud processes and heterogeneous reactions on hydrated aerosols contributed
123 to the formation of oxalic acid (Yang et al., 2009). While the formation mechanism of
124 oxalic acid especially in urban areas is still not clear, online measurements of the
125 mixing state of oxalic acid provide a powerful tool to better understand the formation
126 of oxalic acid in aerosol particles and cloud droplets.

127 The Pearl River Delta (PRD) region has distinct meteorological seasonality
128 under the influence of the Asian monsoon system, which brings air from the ocean in
129 spring and summer, and carries polluted air from northern China in autumn and winter.

130 Strong photochemical activity occurs in summer under the condition of high
131 temperature and relative humidity, and in winter high loadings of particles from
132 northern cities are favorable for the occurrence of haze episode (Bi et al., 2011;Zhang
133 et al., 2013;Zhang et al., 2014). Here we present the seasonal field measurements of
134 the mixing state of oxalic acid-containing particles using a single particle aerosol
135 mass spectrometer (SPAMS) in a rural supersite of the PRD region. The seasonal
136 characteristic of oxalic acid particles and mixing state with secondary species were
137 investigated to explore the formation mechanisms of oxalic acid and aging process of
138 SOA.

139 **2. Methods**

140 **2.1 Aerosol sampling**

141 Particles were sampled using a single particle aerosol mass spectrometer
142 (SPAMS) at the Guangdong Atmospheric Supersite (22.73N, 112.93E), a rural site at
143 Heshan city (Figure S1). The supersite is surrounded by farm land and villages, with
144 no local industrial or traffic emissions. Ambient aerosols were sampled to the SPAMS
145 through a 2.5m long copper tube with 0.5m of the sampling inlet located above the top
146 of the building. The measurement period was from July 18 to August 1 in 2014, and
147 from January 27 to February 8 in 2015. Real-time PM_{2.5} mass concentration was
148 simultaneously measured by a TEOM monitor (series 1405, Thermo scientific), and
149 hourly concentrations of O₃ were measured by an O₃ analyzer (model 49i, Thermo
150 scientific). The local meteorological data including temperature, relative humidity and
151 visibility were measured on the rooftop of the building. The average temperature
152 during the field study was 29.5 °C in summer and 14.1 °C in winter and the average
153 relative humidity was 71.7% and 63% in summer and winter, respectively.

154 **2.2 SPAMS**

155 Real-time measurements of single atmospheric particles has been demonstrated
156 by Prather and co-workers in the 1990s using aerosol time-of-flight mass
157 spectrometry (ATOFMS) (Prather et al., 1994;Noble and Prather, 1996). Based on the
158 same principle, the single particle aerosol mass spectrometer (SPAMS) developed by

159 Guangzhou Hexin Analytical Company was applied to field measurements of single
160 particles in the current work. The details of the SPAMS system have been introduced
161 previously (Li et al., 2011). Briefly, aerosol particles are sampled into the vacuum
162 pumped aerodynamic lens of the SPAMS through an electro-spark machined 80 μ m
163 critical orifice at a flow rate of 75 ml min⁻¹. The individual particles with a terminal
164 velocity are introduced to the sizing region. The velocity of each single particle is
165 detected by two continuous laser beams (diode Nd:YAG, 532 nm) with a space of 6
166 cm. The velocity is then used to calculate the single particle aerodynamic diameter
167 and provide the precise timing of the firing of a 266 nm laser used to induce
168 desorption and ionization (Nd:YAG laser, 266nm). The energy of the
169 desorption/ionization 266 nm laser was 0.6 mJ and the power density was kept at
170 about 1.6 $\times 10^8$ W/cm² during both sampling periods. The 266 nm laser generates
171 positive and negative ions that are detected by a Z-shaped bipolar time of flight mass
172 spectrometer. The size range of the detected single particles is 0.2 to 2 μ m.
173 Polystyrene latex spheres (Nanosphere size standards, Duke Scientific Corp., Palo
174 Alto) of 0.22-2.0 μ m diameter were used for size calibration.

175 **2.3 Data analysis**

176 The size and chemical composition of single particles detected by SPAMS were
177 analyzed using the COCO toolkit based on the Matlab software. Particles were
178 clustered into several groups using the neural network algorithm (ART-2a) to group
179 particles into clusters with similar mass spectrum features. The ART-2a parameters
180 used in this work were set to a vigilance factor of 0.8, a learning rate of 0.05, and a
181 maximum of 20 iterations. We collected 516,679 and 767,986 particles with both
182 positive and negative mass spectra in summer and winter respectively. A standard
183 solution of oxalic acid was prepared with pure oxalic acid (H₂C₂O₄, purity: 99.99%,
184 Aladdin Industrial Corporation) and atomized to aerosols. After drying through two
185 silica gel diffusion driers, pure oxalic acid particles were directly introduced into the
186 SPAMS. The positive and negative mass spectra of oxalic acid are shown in Figure S2.
187 Based on the mass spectra of pure oxalic acid and previous ambient measurements by
188 ATOFMS (Silva and Prather, 2000; Sullivan and Prather, 2007a; Yang et al., 2009),

189 HC_2O_4^- (m/z -89) is selected as the ion peak for oxalic acid containing particles. In
190 this work, oxalic acid particles are identified if the peak area of m/z -89 was larger
191 than 0.5% of the total signal in the mass spectrum. With this threshold, 13109 and
192 20504 of oxalic acid-containing particles were obtained in summer and winter
193 separately, accounting for 2.5% and 2.7% of the total detected particles. The
194 percentage of oxalic acid-containing particles in total particles in this work was
195 comparable to the reported value in the urban area of Shanghai (3.4%) (Yang et al.,
196 2009). However, these percentages are in general much lower than those reported in
197 cleaner environments such as the western Pacific Ocean where oxalic acid was found
198 in up to 1-40% of total particles due to little anthropogenic influences (Sullivan and
199 Prather, 2007a).

200 The oxalic acid containing particles are classified into eight types in the
201 following order: elemental carbon (EC), organic carbon (OC), elemental and organic
202 carbon (ECOC), biomass burning (BB), heavy metal (HM), secondary (Sec),
203 sodium-potassium (NaK) and dust. Different type particles are identified according to
204 characteristic ion markers and dominant chemical species (Table S1): (1) particles
205 containing abundant carbon clusters like $\pm 12[\text{C}]^{+/-}$, $\pm 24[\text{C}_2]^{+/-}$, $\pm 36[\text{C}_3]^{+/-}$ with relative
206 peak area more than 0.5% are classified as EC type, (2) any remaining particles
207 containing abundant signals of $27[\text{C}_2\text{H}_3]^+$, $43[\text{C}_2\text{H}_3\text{O}]^+$ and hydrocarbon clusters with
208 relative peak area more than 0.5% are classified as OC type, (3) any remaining
209 particles containing signals of $\pm 12[\text{C}]^{+/-}$, $\pm 24[\text{C}_2]^{+/-}$, $37[\text{C}_3\text{H}]^+$ and $43[\text{C}_2\text{H}_3\text{O}]^+$ with
210 relative peak area more than 0.5% are classified as ECOC type, (4) any remaining
211 particles containing abundant signals of $39[\text{K}]^+$ (peak area>1500) with relative peak
212 area of $-59[\text{C}_2\text{H}_3\text{O}_2]^-$ and $-73[\text{C}_3\text{H}_5\text{O}_2]^-$ simultaneously more than 0.5% are classified
213 as BB type, (5) any remaining particles containing signals of $55[\text{Mn}]^+$, $56[\text{Fe}]^+$,
214 $63/65[\text{Cu}]^+$, $64[\text{Zn}]^+$ and $208[\text{Pb}]^+$ with relative peak area more than 0.5% are
215 classified as HM type, (6) any remaining particles containing abundant signals of
216 $18[\text{NH}_4]^+$ (peak area>50), $-62[\text{NO}_3]^-$ (peak area>100) and $-97[\text{HSO}_4]^-$ (peak area>100)
217 are classified as Sec type, (7) any remaining particles containing abundant signals of
218 $23[\text{Na}]^+$ (peak area>1500) and related species are classified as NaK type, (8) any

219 remaining particles containing signals of $40[\text{Ca}]^+$, $56[\text{CaO}]^+$ and related species are
220 classified as dust type. The rules for oxalic acid particles classification in the current
221 work have been reported in previous studies (Sullivan and Prather, 2007b; Yang et al.,
222 2009; Zhang et al., 2013; Li et al., 2014).

223 **2.4 Inorganic ions and in-situ pH (pH_{is})**

224 Water-soluble inorganic ions and trace gases were determined by an online
225 analyzer for monitoring aerosols and gases (MARGA, model ADI 2080, Applikon
226 Analytical B. V. Corp., the Netherlands) with a $\text{PM}_{2.5}$ sampling inlet at one hour
227 resolution from July 18 to August 1 in 2014. The principle and instrumental design
228 has been described in detail elsewhere (ten Brink et al., 2007; Du et al., 2011; Behera et
229 al., 2013; Khezri et al., 2013). Standard solutions containing all detected ions were
230 injected into MARGA before and after the field measurement. The liquid water
231 content and the concentration of H^+ in particles are calculated using the ISORROPIA
232 II model (Nenes et al., 1998, 1999; Fountoukis and Nenes, 2007). We choose stable
233 mode and reverse type in the ISORROPIA model to calculate the concentration of H^+
234 and the liquid water content in this work. The in-situ pH (pH_{is}) of particles is
235 calculated through the following equation:

$$236 \quad pH_{is} = -\log \alpha_{\text{H}^+} = -\log(\gamma_{\text{H}^+} \times n_{\text{H}^+} \times 1000/V_a) \quad (1)$$

237 where n_{H^+} is the concentration of H^+ (mol m^{-3}) and V_a is the volume concentration of
238 the H_2O ($\text{cm}^3 \text{m}^{-3}$), while γ_{H^+} is the activity coefficient of H^+ (Xue et al., 2011; Cheng
239 et al., 2015). The temporal variation of pH_{is} of ambient $\text{PM}_{2.5}$ particles is presented in
240 Figure S3, and demonstrated that 97% of particles were acidic in summer.

241 **3. Results and Discussion**

242 **3.1 Seasonal variation of oxalic acid containing particles**

243 The clustered 48 hr back trajectories of air masses arriving in Heshan during the
244 sampling period are shown in Figure S4. In summer, air masses at 500m levels above
245 the ground were mainly from the ocean and rural areas with less influence of human
246 activity, while in winter air masses were directly from urban areas of Guangzhou and

247 Foshan, indicating a strong influence from anthropogenic emissions. The temporal
248 variations of the total detected particles and oxalic acid containing particles in
249 summer and winter are shown in Figure 1. The total particles had similar trends with
250 the mass concentration of ambient PM_{2.5}, suggesting that the counts of total particles
251 detected by SPAMS can be representative of PM_{2.5} mass concentration during the
252 whole sampling periods. The oxalic acid (C₂-containing) particles, in general,
253 exhibited distinct diurnal peaks from July 28 to August 1, while they showed different
254 temporal trends in winter. The relative abundance of oxalic acid particles in all of the
255 sampled particles (C₂/total ratio) had the same variation with the abundance of oxalic
256 acid particles in summer, especially in the period of July 28 – August 1 (Figure 1). In
257 winter, however, particle counts and relative abundance of oxalic acid had different
258 temporal changes except Jan 30 and February 5-8, when the count and relative
259 abundance of oxalic acid particles simultaneously had a sudden increase.

260 The oxalic acid-containing particles were clustered into eight groups, and they
261 altogether accounted for 89.6% and 95.1% of total oxalic acid particles in summer and
262 winter, respectively. Table 1 shows that in summer heavy metal (HM) type particles
263 contributed 31.3% to total oxalic acid particles, followed by the Sec (19.2%) and BB
264 type (13%). However, in winter BB type particles were the most abundant and
265 accounted for 24.2% of the oxalic acid-containing particles, followed by EC and HM
266 type. Besides, carbonaceous type particles including EC, OC, ECOC and BB
267 accounted for 28.1% of oxalic acid particles in summer and 59.8% in winter,
268 indicating the seasonal different characteristics of oxalic acid particles. The temporal
269 variations of eight groups of oxalic acid particles in summer and winter are illustrated
270 in Figure 1. In summer HM type particles (orange color) and total oxalic acid particles
271 exhibited similar diurnal patterns, suggesting a possible connection between the
272 production of oxalic acid and the transition metals (e.g. Fe, Cu) (Zhou et al., 2015).
273 Although Sec, BB and EC type particles showed similar diurnal patterns with total
274 oxalic acid particles, the concentrations of these type particles were generally lower
275 than HM type particles. In winter diurnal variation of oxalic acid particles was not
276 obvious but a sharp increase, accompanied by the increase of BB, EC and Sec type

277 particles, was observed on February 8.

278 The averaged positive and negative ion mass spectra of oxalic acid containing
279 particles are shown in Figure 2. The positive ion spectrum of oxalic acid particles in
280 summer was characterized by high fractions of metal ion peaks including $23[\text{Na}]^+$,
281 $27[\text{Al}]^+$, $39[\text{K}]^+$, $55[\text{Mn}]^+$, $56[\text{Fe}]^+$, $63/65[\text{Cu}]^+$, $64[\text{Zn}]^+$ and $208[\text{Pb}]^+$, and
282 carbonaceous marker ions at m/z $27[\text{C}_2\text{H}_3]^+$, $36[\text{C}_3]^+$, $43[\text{C}_2\text{H}_3\text{O}/\text{C}_3\text{H}_7]^+$, $48[\text{C}_4]^+$
283 (Figure 2 a). The negative ion spectrum of oxalic acid particles in summer was
284 characterized by the strong intensity of secondary ions including m/z $-46[\text{NO}_2]^-$,
285 $-62[\text{NO}_3]^-$, $-79[\text{PO}_3]^-$, $-80[\text{SO}_3]^-$, $-96[\text{SO}_4]^-$ and $-97[\text{HSO}_4]^-$, as well as carbon clusters
286 of $-24[\text{C}_2]^-$, $-36[\text{C}_3]^-$, $-48[\text{C}_4]^-$ and BB markers of $-59[\text{C}_2\text{H}_3\text{O}_2]^-$ and $-73[\text{C}_3\text{H}_5\text{O}_2]^-$
287 (Figure 2 b) (Zauscher et al., 2013). More carbonaceous clusters, i.e., $27[\text{C}_2\text{H}_3]^+$,
288 $29[\text{C}_2\text{H}_5]^+$, $36[\text{C}_3]^+$, $37[\text{C}_3\text{H}]^+$, $43[\text{C}_2\text{H}_3\text{O}]^+$, $48[\text{C}_4]^+$, $51[\text{C}_4\text{H}_3]^+$, $55[\text{C}_4\text{H}_7]^+$, $60[\text{C}_5]^+$,
289 $63[\text{C}_5\text{H}_3]^+$, $65[\text{C}_5\text{H}_5]^+$, $74[\text{C}_2\text{H}_2\text{O}_3]^+$, $77[\text{C}_6\text{H}_5]^+$, were observed in the positive ion
290 spectrum of oxalic acid particles in winter (Figure 2 c) than in summer. The negative
291 ion spectrum of oxalic acid particles in winter (Figure 2 d) contained a large amount
292 of secondary ions, similar to those found in summer, and a more intense signal of
293 nitric acid ($-125[\text{HNO}_3\text{NO}_3]^-$), suggesting an acidic nature of oxalic acid particles in
294 winter.

295 The mixing state of oxalic acid particles with sulfate, nitrate and ammonium
296 (SNA) was investigated through the percentage of SNA-containing oxalic acid
297 particles in total oxalic acid particles (Figure 3). Oxalic acid was found to be
298 internally mixed with sulfate and nitrate during both sampling periods with percentage
299 of 93% and 94% in summer respectively, and both 98% in winter (Figure 3 a).
300 However, the NH_4^+ -containing oxalic acid particle ($\text{C}_2\text{-NH}_4^+$) only accounted for 18%
301 of total oxalic acid particles in summer but this fraction increased to 71% in winter,
302 and linear correlation between $\text{C}_2\text{-NH}_4^+$ particles and total oxalic acid particles
303 showed better linear regression ($r^2=0.98$) in winter than summer, indicating a general
304 mixing state of NH_4^+ with oxalic acid in winter. Aqueous phase production of SO_4^{2-}
305 has been studied well and the linear correlation between oxalic acid and SO_4^{2-} has
306 been used to study the production of oxalic acid through aqueous phase reactions (Yu

307 et al., 2005;Miyazaki et al., 2009;Cheng et al., 2015). In our work, oxalic acid and
308 $C_2-SO_4^{2-}$ displayed good correlations in summer and winter (both $r^2=0.99$), which
309 suggests a common production route of oxalic acid and sulfate, likely aqueous phase
310 reactions.

311 Figure 4 shows the unscaled size-resolved number distributions of the eight types
312 of oxalic acid particles. Oxalic acid mainly existed in 0.4 to 1.2 μm particles during
313 the entire sampling period but exhibited different peak modes for each particle type in
314 summer and winter. In summer, major types of oxalic acid particles showed distinct
315 peak mode at different size diameter. EC and Sec type particles peaked at 0.5 μm ,
316 followed by BB type particles at 0.55 μm , then HM type particles at 0.6 μm , and OC
317 type particles at 0.7 μm . The difference of peak mode suggests the possible different
318 chemical evolution process for each type oxalic acid-containing particles. However, in
319 winter, oxalic acid particles showed broader size distribution from 0.5 to 0.8 μm for
320 all particle types. Oxalic acid particles of all types were generally larger in winter than
321 summer, possibly due to condensation and coagulation of particles during aging of
322 oxalic acid particles in winter.

323 **3.2 Photochemical production of oxalic acid in summer**

324 In summer oxalic acid particles showed peaks in the afternoon especially from
325 July 28 to August 1, which was in agreement with the variation pattern of the O_3
326 concentration (Figure 5), indicating a strong association of oxalic acid formation with
327 photochemical reactions. Malonic acid is another product of photochemical oxidation
328 of organic compounds (Kawamura and Ikushima, 1993;Wang et al., 2012;Meng et al.,
329 2013;Meng et al., 2014). In our campaign, malonic acid containing particles had
330 diurnal trends similar to oxalic acid particles and O_3 concentration. As the dominant
331 particle type, HM particles had identical variation pattern with total oxalic acid
332 particles. They are characterized by highly abundant metal ion peaks like $55[Mn]^+$,
333 $56[Fe]^+$, $63/65[Cu]^+$, $64[Zn]^+$ and $208[Pb]^+$, as well as secondary ion peaks of
334 $-46[NO_2]^-$, $-62[NO_3]^-$, $-80[SO_3]^-$, $-96[SO_4]^-$ and $-97[HSO_4]^-$ in the negative spectrum
335 in summer (Figure 6). In order to investigate the photochemical formation of oxalic
336 acid in summer, the diurnal variations of O_3 , oxalic acid particles, HM type particles

337 and pH_{is} of ambient particles averaged from July 28 to August 1, 2014 are shown in
338 Figure 7. The concentration of O_3 increased after 9:00 and peaked at 17:00, while
339 oxalic acid particles and HM type particles both increased after 10:00 and showed two
340 peaks at 15:00 and 19:00. The pH_{is} of ambient particles ranging from -1.42 to 4.01
341 indicated an acidic environment, and the temporal trends of RH, inorganic ions and
342 H^+ (aq) in aerosols are shown in Figure S5. The oxidation of glyoxal and glyoxylic
343 acid by $\bullet OH$ has been identified as an important pathway of oxalic acid production by
344 field and laboratory studies (Ervens et al., 2004; Ervens and Volkamer, 2010; Wang et
345 al., 2012; Wang et al., 2015). In summer strong photochemical activity and high O_3
346 concentrations in the afternoon lead to more production of dicarbonyls and aldehydes
347 (e.g. glyoxal and methylglyoxal) from VOCs (Myriokefalitakis et al., 2011), which
348 increases the precursors of oxalic acid. The aqueous phase oxidation of glyoxal can
349 take place in both clouds and wet aerosols (Lim et al., 2010). However, the lower
350 yield of oxalic acid from glyoxal in wet aerosols compared to in clouds has been
351 reported in previous chamber experiments due to the formation of substantial amount
352 of high molecular weight products such as oligomers in aerosol-related concentrations
353 (Carlton et al., 2007; Tan et al., 2009). These findings may explain the lower peak of
354 oxalic acid particles at 15:00 compared to that at 19:00. Besides, the precursors of
355 oxalic acid such as glyoxylic acid have higher reaction rate with $\bullet OH$ in high pH
356 solutions according to previous studies (Ervens et al., 2003; Herrmann, 2003; Cheng et
357 al., 2015), and in this work the increase of pH_{is} was observed as the enhancement of
358 oxalic acid particles in the afternoon (Figure 7), which suggests an efficient oxalic
359 acid production from the oxidation of precursors.

360 The similar photochemical pattern of HM type particles with O_3 and total oxalic
361 acid particles implies a possible participation of metal ions in the formation process of
362 oxalic acid. The modeling studies from Ervens et al. (2014) suggest that oxalic acid
363 production from glyoxal and glyoxylic acid in aqueous phase significantly depends on
364 $\bullet OH$ availability (Ervens et al., 2014). The main sources of aqueous phase $\bullet OH$ in
365 cloud droplets include direct uptake from the gas phase (Jacob, 1986), ozone
366 photolysis by UV and visible light at the air-water interface (Anglada et al., 2014),

367 and also aqueous phase chemical reactions(Gligorovski et al., 2015). For the last kind
368 of source, •OH radicals could be generated through Fenton or Fenton like reactions
369 and photolysis of H₂O₂, NO₃⁻, NO₂⁻, and chromophoric dissolved organic matter
370 (CDOM) (Badali et al., 2015;Ervens, 2015;Herrmann et al., 2015;Tong et al., 2016).
371 Given that SPAMS cannot be used to quantify the concentrations of iron ions and
372 H₂O₂, we will investigate the relative contribution of different source •OH radicals to
373 the formation of oxalic acid and show results in our follow up studies.

374 The oxalic acid loss through the photolysis of iron oxalato complexes is a
375 significant sink according to field measurements and model simulations (Sorooshian
376 et al., 2013;Weller et al., 2014;Zhou et al., 2015). Considering the high abundance of
377 iron in oxalic acid particles in the current work (Figure 6), the photolysis of iron
378 oxalato complexes could have played an important role in the diurnal variation of
379 oxalic acid particles. Because the mass concentration of Fe (III) and oxalic acid could
380 not be obtained through SPAMS, the diurnal variations of peak area of iron (m/z=56)
381 and oxalic acid (m/z=-89) were used to investigate the role of iron on the net
382 production of oxalic in the HM type particles from July 28 to August 1, 2014 (Figure
383 8). Interestingly, the peak area of iron exhibited opposite trend with the peak area of
384 oxalic acid from 4:00 to 11:00. As the peak area of Fe increased from 1565 to 29920
385 from 4:00 to 7:00, the peak area of oxalic decreased from 6052 to 3487 accordingly.
386 From 8:00 to 11:00, the peak area of Fe had a very low value of 1168, but the peak
387 area of oxalic had a very high value of 5538. In addition, the peak area of iron
388 exhibited a high value of 138199 at 14:00, while the peak area of oxalic acid showed a
389 lower peak of 7687 at 14:00 and a higher peak of 11879 at 19:00 with an extremely
390 low abundance of iron. Above opposite variation patterns of iron and oxalic acid in
391 iron rich HM type particles during the photochemical activity period from 5:00 to
392 19:00 strongly indicated that photolysis of iron oxalato complexes could be an
393 efficient sink of oxalic acid.

394 The influence from traffic emission was investigated through the diurnal
395 variations of total EC type particles and NO₂ (Figure S6). The EC type particles
396 increased from 12:00 to 21:00, which had same variation as total oxalic acid, but NO₂

397 followed the rush hour pattern with two peaks from 5:00 to 8:00 and from 18:00 to
398 21:00. Traffic emission is not expected to have a large contribution to oxalic acid in
399 this study. The wind speed was low during the whole day (Figure S6), especially
400 between 9:00 and 18:00, which provided a stagnant environment for the increase in
401 oxalic acid produced from photochemical process.

402 **3.3 Formation process of oxalic acid in winter**

403 Despite lower O₃ concentrations and photochemical activity in winter, oxalic
404 acid was still prevalent in carbonaceous particles, especially BB type particles. While
405 oxalic acid was found to be internally mixed with sulfate and nitrate both in summer
406 and winter, the nitric acid was only observed in oxalic acid particles in winter,
407 indicating a strongly acidic nature of oxalic acid particles in winter. Considering a
408 possible connection of oxalic acid production with the acidic environment, the
409 temporal concentrations of oxalic acid, sulfate and nitrate were investigated through
410 their peak areas in the carbonaceous type oxalic acid particles including EC, OC,
411 ECOC and BB type in Figure 9. The peaks of m/z -62[NO₃]⁻ and -97[HSO₄]⁻
412 represent nitrate and sulfate, respectively. Nitrate, sulfate and oxalic acid showed very
413 similar variation patterns in winter, suggesting a close connection of the formation of
414 oxalic acid with the existence of nitrate and sulfate. Although nitric acid was found in
415 the oxalic acid particles, the acidity of the oxalic acid particles was not estimated
416 since the real-time concentration of inorganic ions was not available during the
417 sampling period in winter. Instead the relative acidity ratio (R_{ra}), defined as the ratio
418 of total peak areas of nitrate and sulfate to the peak area of ammonium (m/z
419 18[NH₄]⁺), was used (Denkenberger et al., 2007;Pratt et al., 2009). The R_{ra} of
420 carbonaceous type oxalic acid particles ranged from 7 to 114 with an average value of
421 25 (Figure 9), indicating an intensely acidic environment of carbonaceous type oxalic
422 acid particles in winter. Several studies have reported the formation of oxalic acid
423 through the oxidation of glyoxal and related precursors in acidic aqueous phase
424 (Carlton et al., 2006;Carlton et al., 2007;Tan et al., 2009). Although the influence of
425 different particle acidity on the oxidation process of glyoxal still needs evaluation, the
426 moderate acidic environment is favorable for the production of oxalic acid from the

427 oxidation of glyoxal (Herrmann, 2003;Ervens and Volkamer, 2010;Eugene et al.,
428 2016). In this work the acidic environment of the carbonaceous type oxalic acid
429 particles and similar variation patterns among oxalic acid, sulfate and nitrate may
430 suggest a relationship between the degradation of organic precursors and the acidic
431 chemical process. However, the temporal change of R_{ra} did not follow a similar trend
432 as the peak area of oxalic acid in most particles, possibly due to the multi-step
433 formation of oxalic acid influenced by many factors such as precursors, liquid water
434 content and ion strength (Carlton et al., 2007;Cheng et al., 2013;Cheng et al., 2015).

435 The sharp increase of oxalic acid particles on February 8, 2015 (Figure 1) was
436 selected as a typical episode to investigate the formation processes of oxalic acid in
437 winter. During the episode, the 48 hr back trajectory analysis showed air masses that
438 originated from the urban areas of Guangzhou and Foshan city (Figure S4), indicating
439 strong influence on organic precursors from anthropogenic emissions. Oxalic acid
440 particle types were dominated by BB (23.2%), followed by EC (22.0%) and Sec
441 (15.1%) type (Table 2). Carbonaceous particles including EC, ECOC, OC, BB
442 accounted for 61.6% of the total oxalic acid particles. The mass spectra of oxalic acid
443 particles were characterized by many hydrocarbon clusters of $27[C_2H_3]^+$, $29[C_2H_5]^+$,
444 $37[C_3H]^+$, $43[C_2H_3O]^+$, $51[C_4H_3]^+$, $55[C_4H_7]^+$, $63[C_3H_3]^+$, $65[C_5H_5]^+$, $74[C_2H_2O_3]^+$,
445 $77[C_6H_5]^+$, and carbon clusters of $36[C_3]^+$, $48[C_4]^+$, $60[C_5]^+$ in positive mass spectrum,
446 while the negative mass spectrum was characterized by elemental carbon clusters like
447 $-24[C_2]^-$, $-36[C_3]^-$, $-48[C_4]^-$, biomass burning markers of $-59[C_2H_3O_2]^-$ and
448 $-73[C_3H_5O_2]^-$ and secondary species including $-42[CNO]^-$, $-46[NO_2]^-$, $-62[NO_3]^-$,
449 $-79[PO_3]^-$, $-80[SO_3]^-$, $-96[SO_4]^-$ and $-97[H_2SO_4]^-$ (Figure 10 a).

450 As the precursor of oxalic acid, glyoxal has the potential to react with sulfuric
451 acid to produce organosulfates through acid-catalyzed nucleophilic addition according
452 to laboratory and chamber studies(Surratt et al., 2007;Galloway et al., 2009). The
453 negative ion of $-155([C_2H_3O_2SO_4]^-)$ has been identified as the marker ion of
454 organosulfates derived from glyoxal in chamber and field measurements using
455 ATOFMS (Surratt et al., 2008;Hatch et al., 2011). The formation of organosulfates
456 from glyoxal requires an acidic aqueous environment, which can be used as a marker

457 of acidic aqueous phase aging process of organic compounds. The temporal trend of
458 organosulfate-containing oxalic acid particles in winter is shown in Figure S7, which
459 exhibited a similar pattern as the total oxalic acid particles during the whole sampling
460 period in winter. The percentage of organosulfate-containing oxalic acid particles in
461 total oxalic acid particles ranged from 0 to 16.4% with the highest ratio observed in
462 the episode (February 8). The linear regression between total oxalic acid particles and
463 organosulfate-containing oxalic acid particles in the episode is exhibited in Figure 10b,
464 and the robust correlation ($r^2=0.81$) between them suggests that oxalic acid and
465 organosulfate may share similar formation process. Based on the above discussion,
466 the degradation of carbonaceous species associated with acidic aqueous phase
467 chemical reactions may have an important contribution to the formation of oxalic acid
468 during the episode in winter. Similar particle types and mass spectra of oxalic
469 acid-containing particles during the episode and the whole sampling period in winter
470 were observed, which suggests the acidic aqueous phase chemical processing of
471 organic precursors as a potential source for oxalic acid.

472 **4. Summary and conclusions**

473 Oxalic acid containing particles were measured by a single particle aerosol mass
474 spectrometer (SPAMS) in the summer and winter of 2014 in Heshan, China. They
475 accounted for 2.5% and 2.7% of the total detected ambient particles. In summer heavy
476 metal-containing particles were the largest group of particles containing oxalic acid
477 with a fraction of 31.3% followed by Sec type (19.2%), while in winter BB type was
478 the dominant group with a percentage of 24.2%. More than 90% of oxalic acid
479 particles were internally mixed with sulfate and nitrate during the whole sampling
480 period. Only 18% of oxalic acid particles contained ammonium in summer, which
481 increased to 71% in winter. In summer oxalic acid and O_3 concentration exhibited
482 similar diurnal variations, indicating a substantial contribution of photochemical
483 reactions to oxalic acid formation. The diurnal variations of peak area of iron and
484 oxalic acid in HM type particles indicate a net production of oxalic acid at 15:00
485 lower than at 19:00, likely due to a significant loss of oxalic acid through the

486 photolysis of iron oxalato complexes during the strong photochemical activity period.
487 In winter carbonaceous type particles including EC, OC, ECOC and BB groups
488 accounted for 59.8% of oxalic acid particles and increased to 61.6% in the episode.
489 Nitric acid and organosulfate were found to co-exist in oxalic acid-containing
490 particles in the winter, which suggests a close association with acidic aqueous phase
491 reactions. Acidic aqueous phase chemical processing of organic precursors is a
492 potential contribution for the formation of oxalic acid in winter. The current study
493 demonstrates that SPAMS is a unique tool for understanding the mixing states of
494 different components of ambient aerosols, which are useful for exploring the
495 formation and evolution process of SOA.

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826 **Tables and Figures**

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828 **Table list:**

829

830 Table 1. Summary of major groups of oxalic acid-containing particles in summer and
831 winter in PRD, China.

832

833 Table 2. The abundance of major particle types in total oxalic acid-containing
834 particles during the episode in winter (2/8/2015).

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836 **Figure caption:**

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838 Figure 1. Temporal variations of total detected particles and oxalic acid containing
839 particles during whole sampling periods in Heshan, China: (a) hourly variations of
840 PM_{2.5} mass concentration, total detected particle counts, oxalic acid containing
841 particles, ratio of oxalic acid-containing/total particles and major types of oxalic acid
842 containing particles; (b) variation patterns of relative abundance of major types of
843 oxalic acid containing particles.

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845 Figure 2. The averaged positive and negative ion mass spectra of oxalic acid
846 containing particles is investigated in summer and winter: (a) summer positive, (b)
847 summer negative, (c) winter positive, (d) winter negative. The color bars represent
848 each peak area corresponding to specific fraction in individual particles.

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850 Figure 3. (a) Mixing state of oxalic acid with sulfate, nitrate and ammonium in oxalic
851 acid-containing particles; (b) Linear correlation between NH₄⁺-containing oxalic acid
852 particles and the total oxalic acid particles in summer; (c) Linear correlation between
853 NH₄⁺-containing oxalic acid particles and the total oxalic acid particles in winter.
854 Abbreviations: C₂-NH₄⁺ represents the NH₄⁺-containing oxalic acid particles, and
855 same expressions for C₂-SO₄²⁻ and C₂-NO₃⁻.

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857 Figure 4. Unscaled size-resolved number distributions of major types of oxalic acid
858 particles in summer and winter.

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860 Figure 5. Temporal variations of O₃ concentrations, oxalic acid particles, malonic acid
861 particles and heavy metal type of oxalic acid particles during the entire sampling
862 period in Heshan, China.

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864 Figure 6. The averaged digitized positive and negative ion mass spectra of heavy
865 metal type of oxalic acid-containing particles in summer.

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867 Figure 7. The diurnal variations of O₃ concentration, oxalic acid particles, HM group
868 particles and in-situ pH (pH_{is}) from July 28 to August 1 in 2014.

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870 Figure 8. The diurnal variations of peak area of iron ($m/z=56$) and oxalic acid
871 ($m/z=-89$) in the HM type oxalic acid particles from July 28 to August 1, 2014.

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873 Figure 9. The temporal variations of peak area of nitrate, sulfate and oxalic acid, and
874 the relative acidity ratio (R_{ra}) in carbonaceous type oxalic acid particles in winter.

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876 Figure 10. The comprehensive study of oxalic acid particles increase on Feb 8, 2015:
877 (a) The digitized positive and negative ion mass spectrum of oxalic acid particles
878 during the episode; (b) Linear regression between total oxalic acid particles and
879 organosulfate-containing oxalic acid particles ($m/z -155$).

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Table 1. Summary of major groups of oxalic acid-containing particles in summer and winter in PRD, China.

Particle type	Summer(7/18-8/1, 2014)		Winter(1/27-2/8, 2015)	
	Count	Percentage, %	Count	Percentage, %
EC	1473	11.2	3161	15.4
ECOC	41	0.3	2233	10.9
OC	473	3.6	1922	9.4
BB	1702	13.0	4953	24.2
HM	4104	31.3	3124	15.2
Sec	2511	19.2	2192	10.7
NaK	303	2.3	17	0.1
Dust	1139	8.7	1888	9.2

Abbreviations of major particle types: elemental carbon (EC), elemental and organic carbon (ECOC), organic carbon (OC), biomass burning (BB), heavy metal (HM), secondary (Sec), sodium and potassium (NaK) and dust (Dust).

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Table 2. The abundance of major particle types in total oxalic acid-containing particles during the episode in winter (2/8/2015).

	EC	ECOC	OC	BB	Sec	HM	Dust	other
Count	1250	604	326	1320	856	377	814	132
Percentage, %	22.0	10.6	5.7	23.2	15.1	6.6	14.3	2.3

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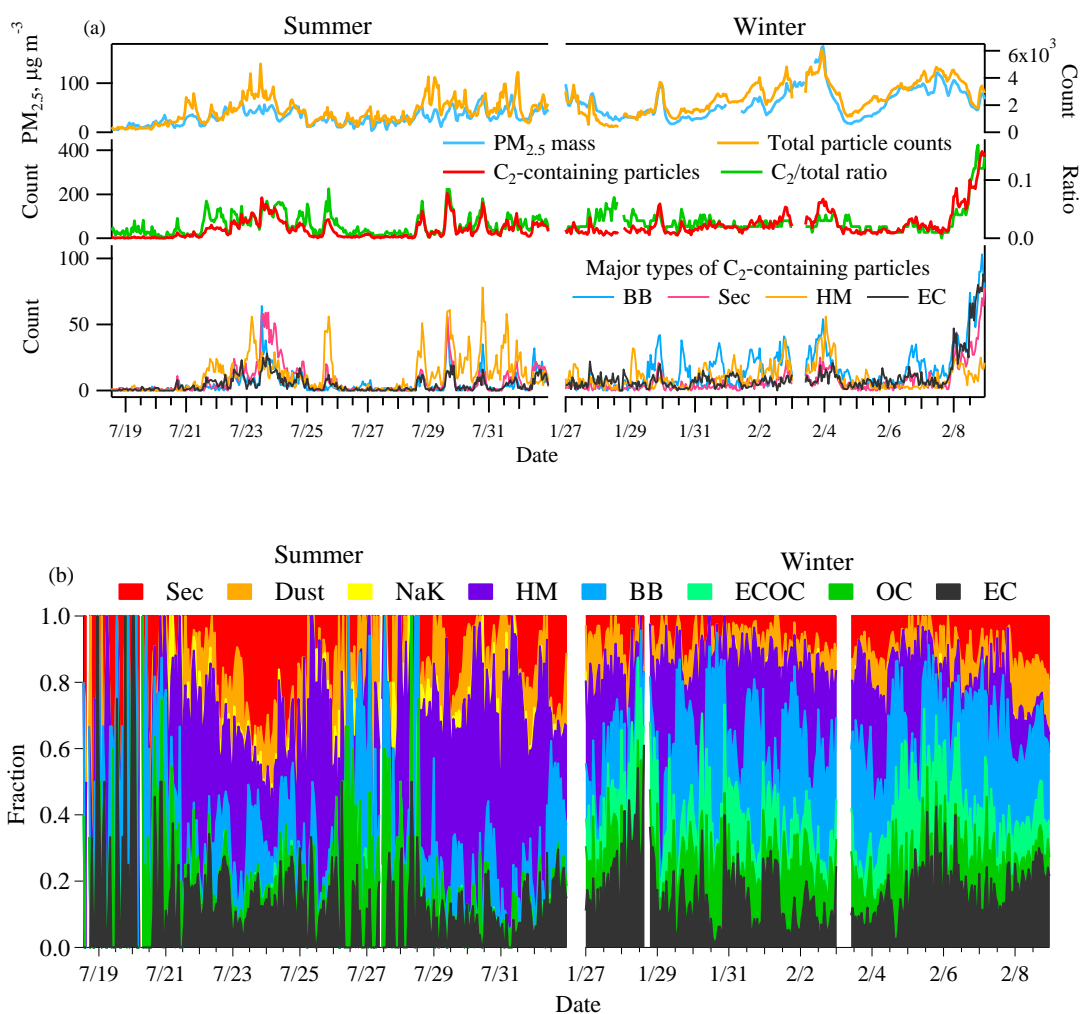
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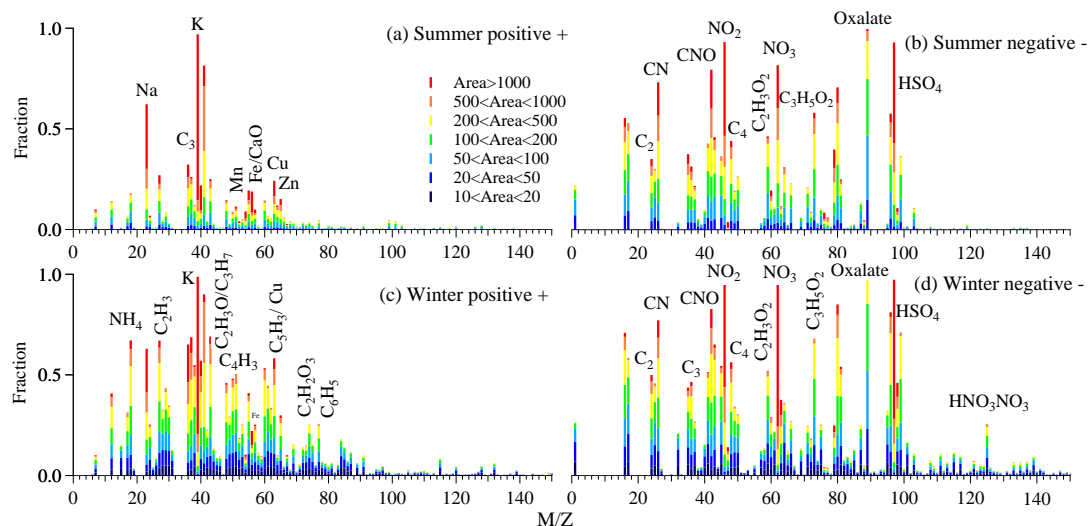


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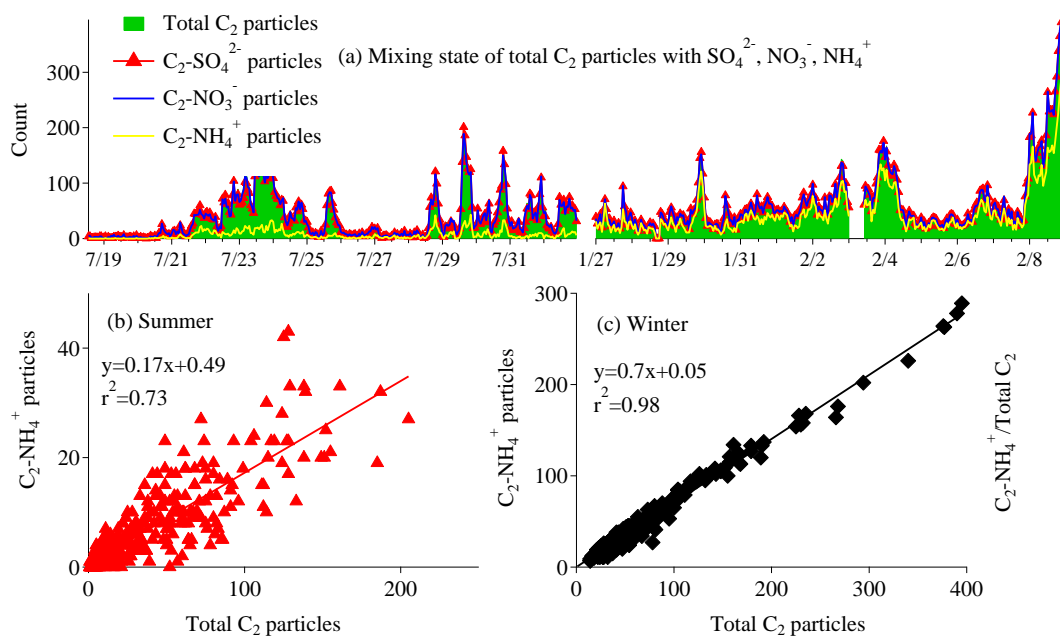
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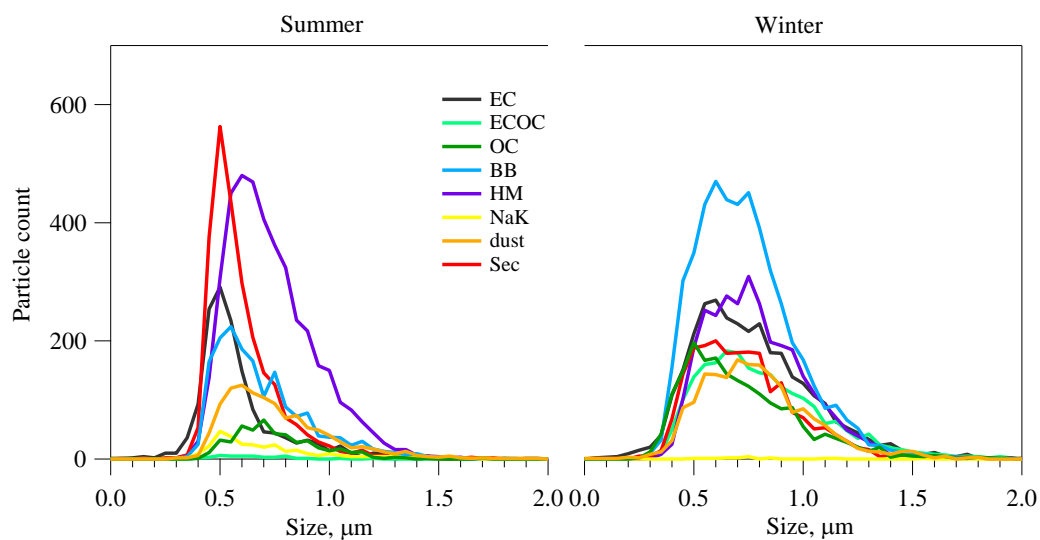
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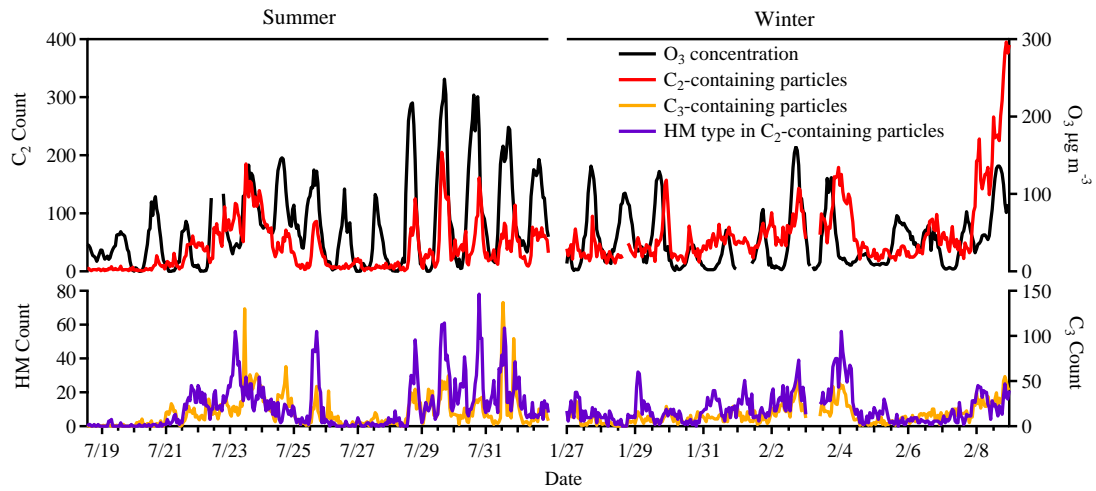


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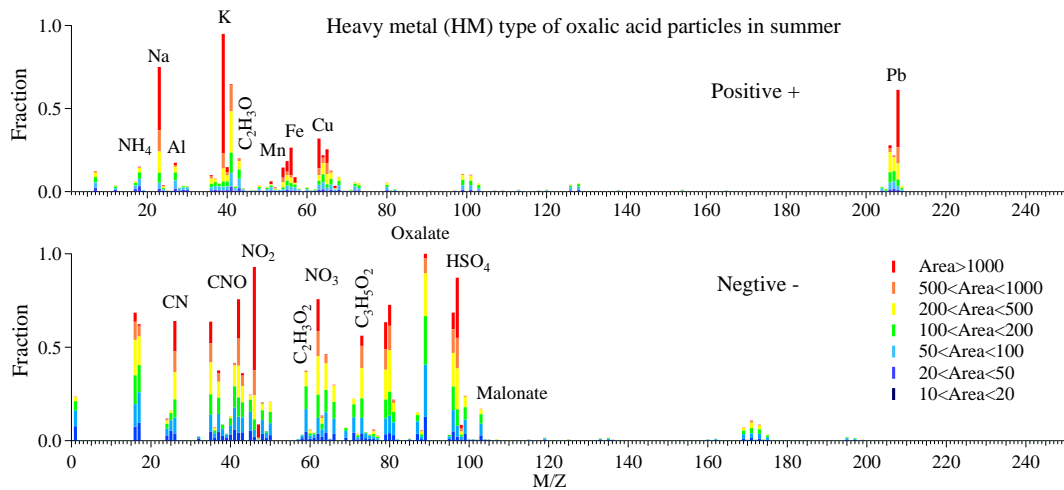
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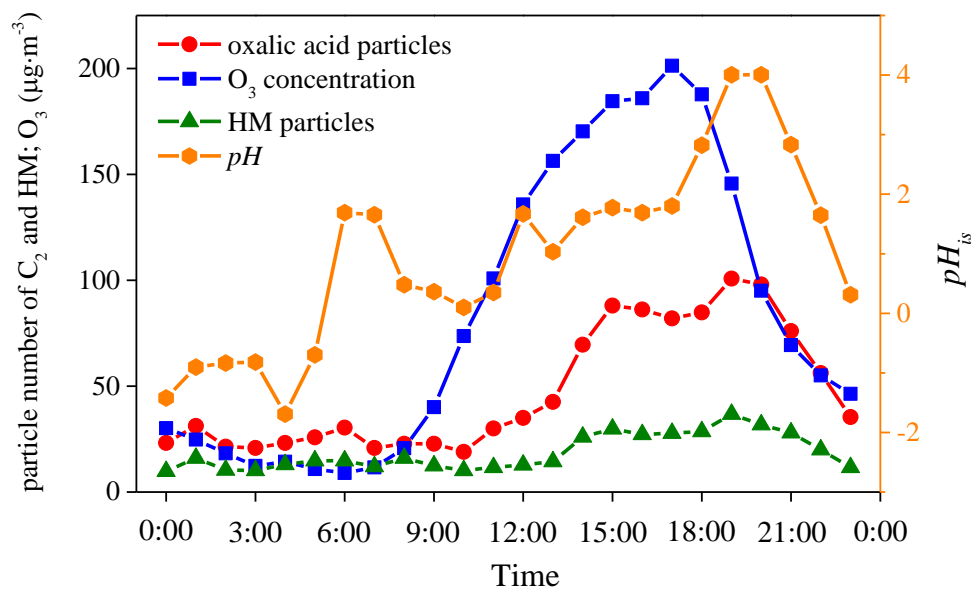
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Figure 5. Temporal variations of O₃ concentrations, oxalic acid particles, malonic acid particles and heavy metal type of oxalic acid particles during the entire sampling period in Heshan, China.



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Figure 6. The averaged digitized positive and negative ion mass spectra of heavy metal type of oxalic acid-containing particles in summer.



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1022 Figure 7. The diurnal variations of O₃ concentration, oxalic acid particles, HM group

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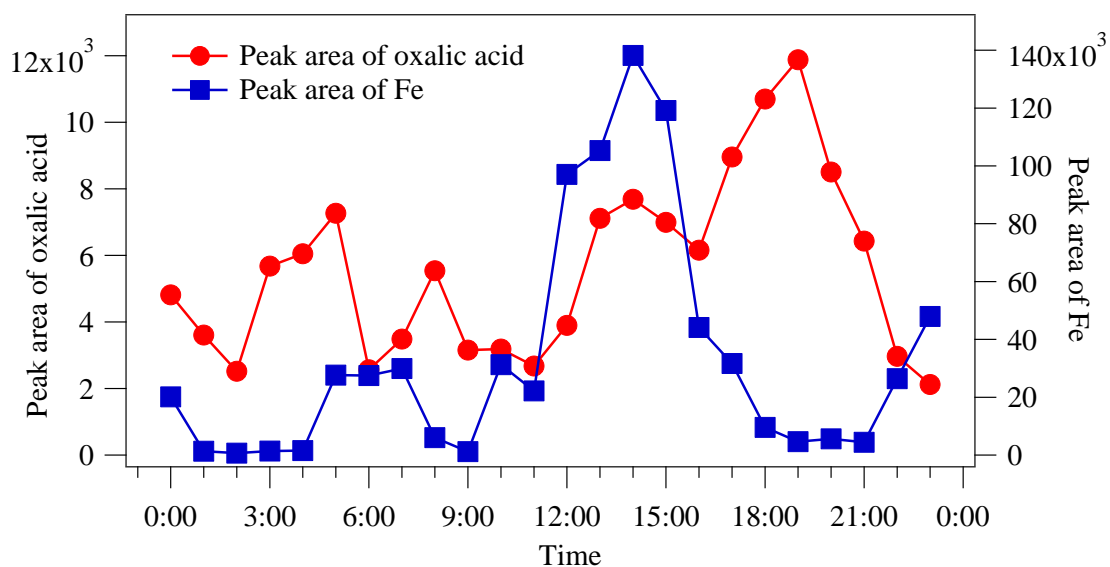
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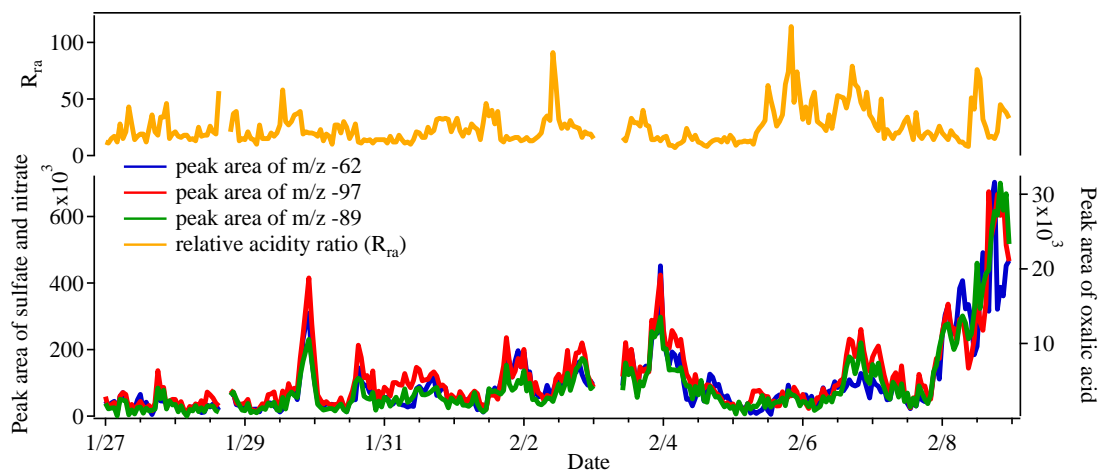
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1030 Figure 8. The diurnal variations of peak area of iron ($m/z=56$) and oxalic acid

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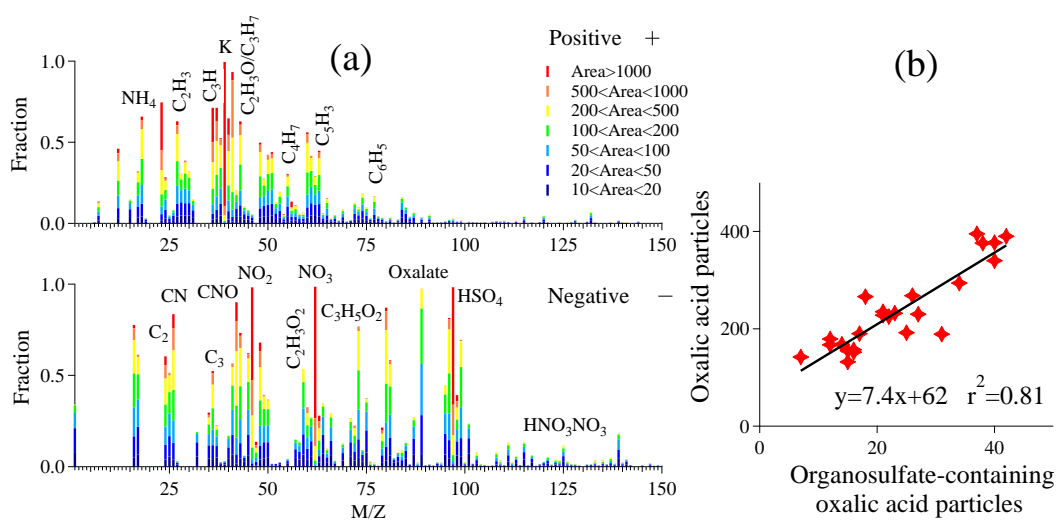
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