

1 Mixing state of oxalic acid containing particles in the rural area of Pearl  
2 River Delta, China: implications for the formation mechanism of oxalic  
3 acid

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42 **Abstract:**

43 The formation of oxalic acid and its mixing state in atmospheric particulate  
44 matter (PM) were studied using a single particle aerosol mass spectrometer (SPAMS)  
45 in the summer and winter of 2014 in Heshan, a supersite in the rural area of the Pearl  
46 River Delta (PRD) region in China. Oxalic acid-containing particles accounted for 2.5%  
47 and 2.7% in total detected ambient particles in summer and winter, respectively.  
48 Oxalic acid was measured in particles classified as elemental carbon (EC), organic  
49 carbon (OC), elemental and organic carbon (ECOC), biomass burning (BB), heavy  
50 metal (HM), secondary (Sec), sodium-potassium (NaK) and dust. Oxalic acid was  
51 found predominantly mixing with sulfate and nitrate during the whole sampling  
52 period, likely due to aqueous phase reactions. In summer, oxalic acid-containing  
53 particle number and ozone concentration followed a very similar trend, which may  
54 reflect the significant contribution of photochemical reactions to oxalic acid formation.  
55 The HM type particles were the most abundant oxalic acid particles in summer and  
56 the diurnal variations of peak area of iron and oxalic acid show opposite trends, which  
57 suggest a possible loss of oxalic acid through the photolysis of iron oxalato complexes  
58 during the strong photochemical activity period. In wintertime, carbonaceous type  
59 particles contained a substantial amount of oxalic acid as well as abundant carbon  
60 clusters and biomass burning markers. The general existence of nitric acid in oxalic  
61 acid-containing particles indicates an acidic environment during the formation process  
62 of oxalic acid. The peak areas of nitrate, sulfate and oxalic had similar temporal  
63 change in the carbonaceous type oxalic acid particles, and the  
64 organosulfate-containing oxalic acid particles well correlated with total oxalic acid  
65 particles during the episode, which suggests the formation of oxalic acid is closely  
66 associated with the oxidation of organic precursors in aqueous phase.

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68 **Keywords:** Oxalic acid; Single particles; Mixing state; Photochemical process;  
69 Aqueous phase reactions.

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## 71 **1. Introduction**

72 Organic aerosol, typically a large fraction of fine particles, contains more than  
73 thousands of organic compounds and contributes to visibility reduction,  
74 photochemical smog, climate change and adverse health effects (Novakov and Penner,  
75 1993;Goldstein and Galbally, 2007;Jimenez et al., 2009;Poschl and Shiraiwa, 2015). A  
76 significant component of organic aerosol is secondary organic aerosol (SOA) formed  
77 from the gas phase oxidation of volatile organic compounds (VOCs) followed by  
78 partitioning of products into particles or from heterogeneous reactions of VOCs with  
79 particles (Hallquist et al., 2009;Zhang et al., 2015). Dicarboxylic acids (DCAs) are  
80 abundant and ubiquitous constituents in SOA and can be effective tracers for the  
81 oxidative processes leading to the formation of SOA (Kawamura and Ikushima,  
82 1993;Ervens et al., 2011;Wang et al., 2012;Cheng et al., 2013). DCAs normally have  
83 high water solubility and low vapor pressure, so they play important roles in  
84 controlling the hygroscopic properties of organic aerosols (Prenni et al., 2003;Ma et  
85 al., 2013) and activating cloud condensation nuclei (Booth et al., 2009). The primary  
86 emissions of DCAs from anthropogenic sources in urban areas are minor (Huang and  
87 Yu, 2007;Stone et al., 2010), and they are mainly derived from secondary oxidation of  
88 VOCs and subsequent intermediates (Ho et al., 2010;Myriokefalitakis et al., 2011).  
89 High concentrations of DCAs have been observed in biomass burning plume (Kundu  
90 et al., 2010;Kawamura et al., 2013) with more than 70% of DCAs produced from  
91 photochemical oxidation of water-soluble organic compounds, and only a small  
92 contribution from direct biomass burning emission (van Pinxteren et al., 2014).

93 The production of DCAs through photochemical reactions has been reported in  
94 many field studies via the analysis of the diurnal and seasonal variations of  
95 DCA(Kawamura and Ikushima, 1993;Kawamura and Yasui, 2005;Aggarwal and  
96 Kawamura, 2008;Pavuluri et al., 2010;Ho et al., 2011;Wang et al., 2017), but the  
97 mechanism of DCAs formation is still not well understood. Oxalic acid is usually the  
98 most abundant DCA observed in the field (Kawamura et al., 2004;Ho et al.,  
99 2007;Kawamura et al., 2010). A number of ground based and airborne field studies

100 have found a tight correlation between oxalic acid and sulfate in ambient particles and  
101 cloud droplets, relating aqueous phase chemistry to the formation of oxalic acid in  
102 aerosols and cloud droplets (Yao et al., 2002; Yao et al., 2003; Yu et al.,  
103 2005; Sorooshian et al., 2006; Sorooshian et al., 2007a; Sorooshian et al.,  
104 2007b; Miyazaki et al., 2009; Wonaschuetz et al., 2012; Wang et al., 2016). In recent  
105 years, several model and laboratory studies suggested that the aqueous phase  
106 oxidation of highly water-soluble organics like glyoxal, methylglyoxal and glyoxylic  
107 acid can efficiently produce oxalic acid in aerosol particles and cloud droplets (Lim et  
108 al., 2010; Myriokefalitakis et al., 2011; Ervens et al., 2014; Yu et al., 2014; McNeill,  
109 2015). Recent stable carbon isotope studies and field observations have also suggested  
110 that oxalic acid forms through aqueous phase reactions (Wang et al., 2012; Cheng et al.,  
111 2015). However, the formation process of oxalic acid in ambient aerosols is still  
112 associated with great uncertainty due to the oxidation rates of precursors and oxidant  
113 levels in photochemistry and aqueous phase chemistry, which needs to be further  
114 studied.

115 Online measurements of the size distribution of oxalic acid-containing particles  
116 and the mixing state of oxalic acid with other compounds in aerosols are useful to  
117 examine the formation and evolution of oxalic acid and SOA particles. Sullivan and  
118 Prather investigated the diurnal cycle and mixing state of DCA-containing particles in  
119 Asian aerosol outflow using aerosol time-of-flight mass spectrometry (ATOFMS), and  
120 proposed the formation of DCA on Asian dust (Sullivan and Prather, 2007a). In  
121 addition, Yang et al. (2009) measured oxalic acid particles in Shanghai and proposed  
122 that in-cloud processes and heterogeneous reactions on hydrated aerosols contributed  
123 to the formation of oxalic acid (Yang et al., 2009). While the formation mechanism of  
124 oxalic acid especially in urban areas is still not clear, online measurements of the  
125 mixing state of oxalic acid provide a powerful tool to better understand the formation  
126 of oxalic acid in aerosol particles and cloud droplets.

127 The Pearl River Delta (PRD) region has distinct meteorological seasonality  
128 under the influence of the Asian monsoon system, which brings air from the ocean in  
129 spring and summer, and carries polluted air from northern China in autumn and winter.

130 Strong photochemical activity occurs in summer under the condition of high  
131 temperature and relative humidity, and in winter high loadings of particles from  
132 northern cities are favorable for the occurrence of haze episode (Bi et al., 2011;Zhang  
133 et al., 2013;Zhang et al., 2014). Here we present the seasonal field measurements of  
134 the mixing state of oxalic acid-containing particles using a single particle aerosol  
135 mass spectrometer (SPAMS) in a rural supersite of the PRD region. The seasonal  
136 characteristic of oxalic acid particles and mixing state with secondary species were  
137 investigated to explore the formation mechanisms of oxalic acid and aging process of  
138 SOA.

## 139 **2. Methods**

### 140 **2.1 Aerosol sampling**

141 Particles were sampled using a single particle aerosol mass spectrometer  
142 (SPAMS) at the Guangdong Atmospheric Supersite (22.73N, 112.93E), a rural site at  
143 Heshan city (Figure S1). The supersite is surrounded by farm land and villages, with  
144 no local industrial or traffic emissions. Ambient aerosols were sampled to the SPAMS  
145 through a 2.5m long copper tube with 0.5m of the sampling inlet located above the top  
146 of the building. The measurement period was from July 18 to August 1 in 2014, and  
147 from January 27 to February 8 in 2015. Real-time PM<sub>2.5</sub> mass concentration was  
148 simultaneously measured by a TEOM monitor (series 1405, Thermo scientific), and  
149 hourly concentrations of O<sub>3</sub> were measured by an O<sub>3</sub> analyzer (model 49i, Thermo  
150 scientific). The local meteorological data including temperature, relative humidity and  
151 visibility were measured on the rooftop of the building. The average temperature  
152 during the field study was 29.5 °C in summer and 14.1 °C in winter and the average  
153 relative humidity was 71.7% and 63% in summer and winter, respectively.

### 154 **2.2 SPAMS**

155 Real-time measurements of single atmospheric particles has been demonstrated  
156 by Prather and co-workers in the 1990s using aerosol time-of-flight mass  
157 spectrometry (ATOFMS) (Prather et al., 1994;Noble and Prather, 1996). Based on the  
158 same principle, the single particle aerosol mass spectrometer (SPAMS) developed by

159 Guangzhou Hexin Analytical Company was applied to field measurements of single  
160 particles in the current work. The details of the SPAMS system have been introduced  
161 previously (Li et al., 2011). Briefly, aerosol particles are sampled into the vacuum  
162 pumped aerodynamic lens of the SPAMS through an electro-spark machined 80 $\mu$ m  
163 critical orifice at a flow rate of 75 ml min<sup>-1</sup>. The individual particles with a terminal  
164 velocity are introduced to the sizing region. The velocity of each single particle is  
165 detected by two continuous laser beams (diode Nd:YAG, 532 nm) with a space of 6  
166 cm. The velocity is then used to calculate the single particle aerodynamic diameter  
167 and provide the precise timing of the firing of a 266 nm laser used to induce  
168 desorption and ionization (Nd:YAG laser, 266nm). The energy of the  
169 desorption/ionization 266 nm laser was 0.6 mJ and the power density was kept at  
170 about 1.6 $\times 10^8$  W/cm<sup>2</sup> during both sampling periods. The 266 nm laser generates  
171 positive and negative ions that are detected by a Z-shaped bipolar time of flight mass  
172 spectrometer. The size range of the detected single particles is 0.2 to 2  $\mu$ m.  
173 Polystyrene latex spheres (Nanosphere size standards, Duke Scientific Corp., Palo  
174 Alto) of 0.22-2.0  $\mu$ m diameter were used for size calibration.

### 175 **2.3 Data analysis**

176 The size and chemical composition of single particles detected by SPAMS were  
177 analyzed using the COCO toolkit based on the Matlab software. Particles were  
178 clustered into several groups using the neural network algorithm (ART-2a) to group  
179 particles into clusters with similar mass spectrum features. The ART-2a parameters  
180 used in this work were set to a vigilance factor of 0.8, a learning rate of 0.05, and a  
181 maximum of 20 iterations. We collected 516,679 and 767,986 particles with both  
182 positive and negative mass spectra in summer and winter respectively. A standard  
183 solution of oxalic acid was prepared with pure oxalic acid (H<sub>2</sub>C<sub>2</sub>O<sub>4</sub>, purity: 99.99%,  
184 Aladdin Industrial Corporation) and atomized to aerosols. After drying through two  
185 silica gel diffusion driers, pure oxalic acid particles were directly introduced into the  
186 SPAMS. The positive and negative mass spectra of oxalic acid are shown in Figure S2.  
187 Based on the mass spectra of pure oxalic acid and previous ambient measurements by  
188 ATOFMS (Silva and Prather, 2000; Sullivan and Prather, 2007a; Yang et al., 2009),

189  $\text{HC}_2\text{O}_4^-$  ( $m/z$  -89) is selected as the ion peak for oxalic acid containing particles. In  
190 this work, oxalic acid particles are identified if the peak area of  $m/z$  -89 was larger  
191 than 0.5% of the total signal in the mass spectrum. With this threshold, 13109 and  
192 20504 of oxalic acid-containing particles were obtained in summer and winter  
193 separately, accounting for 2.5% and 2.7% of the total detected particles. The  
194 percentage of oxalic acid-containing particles in total particles in this work was  
195 comparable to the reported value in the urban area of Shanghai (3.4%) (Yang et al.,  
196 2009). However, these percentages are in general much lower than those reported in  
197 cleaner environments such as the western Pacific Ocean where oxalic acid was found  
198 in up to 1-40% of total particles due to little anthropogenic influences (Sullivan and  
199 Prather, 2007a).

200 The oxalic acid containing particles are classified into eight types in the  
201 following order: elemental carbon (EC), organic carbon (OC), elemental and organic  
202 carbon (ECOC), biomass burning (BB), heavy metal (HM), secondary (Sec),  
203 sodium-potassium (NaK) and dust. Different type particles are identified according to  
204 characteristic ion markers and dominant chemical species (Table S1): (1) particles  
205 containing abundant carbon clusters like  $\pm 12[\text{C}]^{+/-}$ ,  $\pm 24[\text{C}_2]^{+/-}$ ,  $\pm 36[\text{C}_3]^{+/-}$  with relative  
206 peak area more than 0.5% are classified as EC type, (2) any remaining particles  
207 containing abundant signals of  $27[\text{C}_2\text{H}_3]^+$ ,  $43[\text{C}_2\text{H}_3\text{O}]^+$  and hydrocarbon clusters with  
208 relative peak area more than 0.5% are classified as OC type, (3) any remaining  
209 particles containing signals of  $\pm 12[\text{C}]^{+/-}$ ,  $\pm 24[\text{C}_2]^{+/-}$ ,  $37[\text{C}_3\text{H}]^+$  and  $43[\text{C}_2\text{H}_3\text{O}]^+$  with  
210 relative peak area more than 0.5% are classified as ECOC type, (4) any remaining  
211 particles containing abundant signals of  $39[\text{K}]^+$  (peak area>1500) with relative peak  
212 area of  $-59[\text{C}_2\text{H}_3\text{O}_2]^-$  and  $-73[\text{C}_3\text{H}_5\text{O}_2]^-$  simultaneously more than 0.5% are classified  
213 as BB type, (5) any remaining particles containing signals of  $55[\text{Mn}]^+$ ,  $56[\text{Fe}]^+$ ,  
214  $63/65[\text{Cu}]^+$ ,  $64[\text{Zn}]^+$  and  $208[\text{Pb}]^+$  with relative peak area more than 0.5% are  
215 classified as HM type, (6) any remaining particles containing abundant signals of  
216  $18[\text{NH}_4]^+$  (peak area>50),  $-62[\text{NO}_3]^-$  (peak area>100) and  $-97[\text{HSO}_4]^-$  (peak area>100)  
217 are classified as Sec type, (7) any remaining particles containing abundant signals of  
218  $23[\text{Na}]^+$  (peak area>1500) and related species are classified as NaK type, (8) any

219 remaining particles containing signals of  $40[\text{Ca}]^+$ ,  $56[\text{CaO}]^+$  and related species are  
220 classified as dust type. The rules for oxalic acid particles classification in the current  
221 work have been reported in previous studies (Sullivan and Prather, 2007b; Yang et al.,  
222 2009; Zhang et al., 2013; Li et al., 2014).

## 223 **2.4 Inorganic ions and in-situ pH ( $pH_{is}$ )**

224 Water-soluble inorganic ions and trace gases were determined by an online  
225 analyzer for monitoring aerosols and gases (MARGA, model ADI 2080, Applikon  
226 Analytical B. V. Corp., the Netherlands) with a  $\text{PM}_{2.5}$  sampling inlet at one hour  
227 resolution from July 18 to August 1 in 2014. The principle and instrumental design  
228 has been described in detail elsewhere (ten Brink et al., 2007; Du et al., 2011; Behera et  
229 al., 2013; Khezri et al., 2013). Standard solutions containing all detected ions were  
230 injected into MARGA before and after the field measurement. The liquid water  
231 content and the concentration of  $\text{H}^+$  in particles are calculated using the ISORROPIA  
232 II model (Nenes et al., 1998, 1999; Fountoukis and Nenes, 2007). We choose stable  
233 mode and reverse type in the ISORROPIA model to calculate the concentration of  $\text{H}^+$   
234 and the liquid water content in this work. The in-situ pH ( $pH_{is}$ ) of particles is  
235 calculated through the following equation:

$$236 \quad pH_{is} = -\log \alpha_{\text{H}^+} = -\log(\gamma_{\text{H}^+} \times n_{\text{H}^+} \times 1000/V_a) \quad (1)$$

237 where  $n_{\text{H}^+}$  is the concentration of  $\text{H}^+$  ( $\text{mol m}^{-3}$ ) and  $V_a$  is the volume concentration of  
238 the  $\text{H}_2\text{O}$  ( $\text{cm}^3 \text{m}^{-3}$ ), while  $\gamma_{\text{H}^+}$  is the activity coefficient of  $\text{H}^+$  (Xue et al., 2011; Cheng  
239 et al., 2015). The temporal variation of  $pH_{is}$  of ambient  $\text{PM}_{2.5}$  particles is presented in  
240 Figure S3, and demonstrated that 97% of particles were acidic in summer.

## 241 **3. Results and Discussion**

### 242 **3.1 Seasonal variation of oxalic acid containing particles**

243 The clustered 48 hr back trajectories of air masses arriving in Heshan during the  
244 sampling period are shown in Figure S4. In summer, air masses at 500m levels above  
245 the ground were mainly from the ocean and rural areas with less influence of human  
246 activity, while in winter air masses were directly from urban areas of Guangzhou and



247 Foshan, indicating a strong influence from anthropogenic emissions. The temporal  
248 variations of the total detected particles and oxalic acid containing particles in  
249 summer and winter are shown in Figure 1. The total particles had similar trends with  
250 the mass concentration of ambient PM<sub>2.5</sub>, suggesting that the counts of total particles  
251 detected by SPAMS can be representative of PM<sub>2.5</sub> mass concentration during the  
252 whole sampling periods. The oxalic acid (C<sub>2</sub>-containing) particles, in general,  
253 exhibited distinct diurnal peaks from July 28 to August 1, while they showed different  
254 temporal trends in winter. The relative abundance of oxalic acid particles in all of the  
255 sampled particles (C<sub>2</sub>/total ratio) had the same variation with the abundance of oxalic  
256 acid particles in summer, especially in the period of July 28 – August 1 (Figure 1). In  
257 winter, however, particle counts and relative abundance of oxalic acid had different  
258 temporal changes except Jan 30 and February 5-8, when the count and relative  
259 abundance of oxalic acid particles simultaneously had a sudden increase.

260 The oxalic acid-containing particles were clustered into eight groups, and they  
261 altogether accounted for 89.6% and 95.1% of total oxalic acid particles in summer and  
262 winter, respectively. Table 1 shows that in summer heavy metal (HM) type particles  
263 contributed 31.3% to total oxalic acid particles, followed by the Sec (19.2%) and BB  
264 type (13%). However, in winter BB type particles were the most abundant and  
265 accounted for 24.2% of the oxalic acid-containing particles, followed by EC and HM  
266 type. Besides, carbonaceous type particles including EC, OC, ECOC and BB  
267 accounted for 28.1% of oxalic acid particles in summer and 59.8% in winter,  
268 indicating the seasonal different characteristics of oxalic acid particles. The temporal  
269 variations of eight groups of oxalic acid particles in summer and winter are illustrated  
270 in Figure 1. In summer HM type particles (orange color) and total oxalic acid particles  
271 exhibited similar diurnal patterns, suggesting a possible connection between the  
272 production of oxalic acid and the transition metals (e.g. Fe, Cu) (Zhou et al., 2015).  
273 Although Sec, BB and EC type particles showed similar diurnal patterns with total  
274 oxalic acid particles, the concentrations of these type particles were generally lower  
275 than HM type particles. In winter diurnal variation of oxalic acid particles was not  
276 obvious but a sharp increase, accompanied by the increase of BB, EC and Sec type

277 particles, was observed on February 8.

278 The averaged positive and negative ion mass spectra of oxalic acid containing  
279 particles are shown in Figure 2. The positive ion spectrum of oxalic acid particles in  
280 summer was characterized by high fractions of metal ion peaks including  $23[\text{Na}]^+$ ,  
281  $27[\text{Al}]^+$ ,  $39[\text{K}]^+$ ,  $55[\text{Mn}]^+$ ,  $56[\text{Fe}]^+$ ,  $63/65[\text{Cu}]^+$ ,  $64[\text{Zn}]^+$  and  $208[\text{Pb}]^+$ , and  
282 carbonaceous marker ions at  $m/z$   $27[\text{C}_2\text{H}_3]^+$ ,  $36[\text{C}_3]^+$ ,  $43[\text{C}_2\text{H}_3\text{O}/\text{C}_3\text{H}_7]^+$ ,  $48[\text{C}_4]^+$   
283 (Figure 2 a). The negative ion spectrum of oxalic acid particles in summer was  
284 characterized by the strong intensity of secondary ions including  $m/z$   $-46[\text{NO}_2]^-$ ,  
285  $-62[\text{NO}_3]^-$ ,  $-79[\text{PO}_3]^-$ ,  $-80[\text{SO}_3]^-$ ,  $-96[\text{SO}_4]^-$  and  $-97[\text{HSO}_4]^-$ , as well as carbon clusters  
286 of  $-24[\text{C}_2]^-$ ,  $-36[\text{C}_3]^-$ ,  $-48[\text{C}_4]^-$  and BB markers of  $-59[\text{C}_2\text{H}_3\text{O}_2]^-$  and  $-73[\text{C}_3\text{H}_5\text{O}_2]^-$   
287 (Figure 2 b) (Zauscher et al., 2013). More carbonaceous clusters, i.e.,  $27[\text{C}_2\text{H}_3]^+$ ,  
288  $29[\text{C}_2\text{H}_5]^+$ ,  $36[\text{C}_3]^+$ ,  $37[\text{C}_3\text{H}]^+$ ,  $43[\text{C}_2\text{H}_3\text{O}]^+$ ,  $48[\text{C}_4]^+$ ,  $51[\text{C}_4\text{H}_3]^+$ ,  $55[\text{C}_4\text{H}_7]^+$ ,  $60[\text{C}_5]^+$ ,  
289  $63[\text{C}_5\text{H}_3]^+$ ,  $65[\text{C}_5\text{H}_5]^+$ ,  $74[\text{C}_2\text{H}_2\text{O}_3]^+$ ,  $77[\text{C}_6\text{H}_5]^+$ , were observed in the positive ion  
290 spectrum of oxalic acid particles in winter (Figure 2 c) than in summer. The negative  
291 ion spectrum of oxalic acid particles in winter (Figure 2 d) contained a large amount  
292 of secondary ions, similar to those found in summer, and a more intense signal of  
293 nitric acid ( $-125[\text{HNO}_3\text{NO}_3]^-$ ), suggesting an acidic nature of oxalic acid particles in  
294 winter.

295 The mixing state of oxalic acid particles with sulfate, nitrate and ammonium  
296 (SNA) was investigated through the percentage of SNA-containing oxalic acid  
297 particles in total oxalic acid particles (Figure 3). Oxalic acid was found to be  
298 internally mixed with sulfate and nitrate during both sampling periods with percentage  
299 of 93% and 94% in summer respectively, and both 98% in winter (Figure 3 a).  
300 However, the  $\text{NH}_4^+$ -containing oxalic acid particle ( $\text{C}_2\text{-NH}_4^+$ ) only accounted for 18%  
301 of total oxalic acid particles in summer but this fraction increased to 71% in winter,  
302 and linear correlation between  $\text{C}_2\text{-NH}_4^+$  particles and total oxalic acid particles  
303 showed better linear regression ( $r^2=0.98$ ) in winter than summer, indicating a general  
304 mixing state of  $\text{NH}_4^+$  with oxalic acid in winter. Aqueous phase production of  $\text{SO}_4^{2-}$   
305 has been studied well and the linear correlation between oxalic acid and  $\text{SO}_4^{2-}$  has  
306 been used to study the production of oxalic acid through aqueous phase reactions (Yu

307 et al., 2005;Miyazaki et al., 2009;Cheng et al., 2015). In our work, oxalic acid and  
308  $C_2-SO_4^{2-}$  displayed good correlations in summer and winter (both  $r^2=0.99$ ), which  
309 suggests a common production route of oxalic acid and sulfate, likely aqueous phase  
310 reactions.

311 Figure 4 shows the unscaled size-resolved number distributions of the eight types  
312 of oxalic acid particles. Oxalic acid mainly existed in 0.4 to 1.2  $\mu m$  particles during  
313 the entire sampling period but exhibited different peak modes for each particle type in  
314 summer and winter. In summer, major types of oxalic acid particles showed distinct  
315 peak mode at different size diameter. EC and Sec type particles peaked at 0.5  $\mu m$ ,  
316 followed by BB type particles at 0.55  $\mu m$ , then HM type particles at 0.6  $\mu m$ , and OC  
317 type particles at 0.7  $\mu m$ . The difference of peak mode suggests the possible different  
318 chemical evolution process for each type oxalic acid-containing particles. However, in  
319 winter, oxalic acid particles showed broader size distribution from 0.5 to 0.8  $\mu m$  for  
320 all particle types. Oxalic acid particles of all types were generally larger in winter than  
321 summer, possibly due to condensation and coagulation of particles during aging of  
322 oxalic acid particles in winter.

### 323 **3.2 Photochemical production of oxalic acid in summer**

324 In summer oxalic acid particles showed peaks in the afternoon especially from  
325 July 28 to August 1, which was in agreement with the variation pattern of the  $O_3$   
326 concentration (Figure 5), indicating a strong association of oxalic acid formation with  
327 photochemical reactions. Malonic acid is another product of photochemical oxidation  
328 of organic compounds (Kawamura and Ikushima, 1993;Wang et al., 2012;Meng et al.,  
329 2013;Meng et al., 2014). In our campaign, malonic acid containing particles had  
330 diurnal trends similar to oxalic acid particles and  $O_3$  concentration. As the dominant  
331 particle type, HM particles had identical variation pattern with total oxalic acid  
332 particles. They are characterized by highly abundant metal ion peaks like  $55[Mn]^+$ ,  
333  $56[Fe]^+$ ,  $63/65[Cu]^+$ ,  $64[Zn]^+$  and  $208[Pb]^+$ , as well as secondary ion peaks of  
334  $-46[NO_2]^-$ ,  $-62[NO_3]^-$ ,  $-80[SO_3]^-$ ,  $-96[SO_4]^-$  and  $-97[HSO_4]^-$  in the negative spectrum  
335 in summer (Figure 6). In order to investigate the photochemical formation of oxalic  
336 acid in summer, the diurnal variations of  $O_3$ , oxalic acid particles, HM type particles

337 and  $pH_{is}$  of ambient particles averaged from July 28 to August 1, 2014 are shown in  
338 Figure 7. The concentration of  $O_3$  increased after 9:00 and peaked at 17:00, while  
339 oxalic acid particles and HM type particles both increased after 10:00 and showed two  
340 peaks at 15:00 and 19:00. The  $pH_{is}$  of ambient particles ranging from -1.42 to 4.01  
341 indicated an acidic environment, and the temporal trends of RH, inorganic ions and  
342  $H^+$  (aq) in aerosols are shown in Figure S5. The oxidation of glyoxal and glyoxylic  
343 acid by  $\bullet OH$  has been identified as an important pathway of oxalic acid production by  
344 field and laboratory studies (Ervens et al., 2004; Ervens and Volkamer, 2010; Wang et  
345 al., 2012; Wang et al., 2015). In summer strong photochemical activity and high  $O_3$   
346 concentrations in the afternoon lead to more production of dicarbonyls and aldehydes  
347 (e.g. glyoxal and methylglyoxal) from VOCs (Myriokefalitakis et al., 2011), which  
348 increases the precursors of oxalic acid. The aqueous phase oxidation of glyoxal can  
349 take place in both clouds and wet aerosols (Lim et al., 2010). However, the lower  
350 yield of oxalic acid from glyoxal in wet aerosols compared to in clouds has been  
351 reported in previous chamber experiments due to the formation of substantial amount  
352 of high molecular weight products such as oligomers in aerosol-related concentrations  
353 (Carlton et al., 2007; Tan et al., 2009). These findings may explain the lower peak of  
354 oxalic acid particles at 15:00 compared to that at 19:00. Besides, the precursors of  
355 oxalic acid such as glyoxylic acid have higher reaction rate with  $\bullet OH$  in high pH  
356 solutions according to previous studies (Ervens et al., 2003; Herrmann, 2003; Cheng et  
357 al., 2015), and in this work the increase of  $pH_{is}$  was observed as the enhancement of  
358 oxalic acid particles in the afternoon (Figure 7), which suggests an efficient oxalic  
359 acid production from the oxidation of precursors.

360 The similar photochemical pattern of HM type particles with  $O_3$  and total oxalic  
361 acid particles implies a possible participation of metal ions in the formation process of  
362 oxalic acid. The modeling studies from Ervens et al. (2014) suggest that oxalic acid  
363 production from glyoxal and glyoxylic acid in aqueous phase significantly depends on  
364  $\bullet OH$  availability (Ervens et al., 2014). The main sources of aqueous phase  $\bullet OH$  in  
365 cloud droplets include direct uptake from the gas phase (Jacob, 1986), ozone  
366 photolysis by UV and visible light at the air-water interface (Anglada et al., 2014),

367 and also aqueous phase chemical reactions (Gligorovski et al., 2015). For the last kind  
368 of source, •OH radicals could be generated through Fenton or Fenton like reactions  
369 and photolysis of H<sub>2</sub>O<sub>2</sub>, NO<sub>3</sub><sup>-</sup>, NO<sub>2</sub><sup>-</sup>, and chromophoric dissolved organic matter  
370 (CDOM) (Badali et al., 2015; Ervens, 2015; Herrmann et al., 2015; Tong et al., 2016).  
371 Given that SPAMS cannot be used to quantify the concentrations of iron ions and  
372 H<sub>2</sub>O<sub>2</sub>, we will investigate the relative contribution of different source •OH radicals to  
373 the formation of oxalic acid and show results in our follow up studies.

374 The oxalic acid loss through the photolysis of iron oxalato complexes is a  
375 significant sink according to field measurements and model simulations (Sorooshian  
376 et al., 2013; Weller et al., 2014; Zhou et al., 2015). Considering the high abundance of  
377 iron in oxalic acid particles in the current work (Figure 6), the photolysis of iron  
378 oxalato complexes could have played an important role in the diurnal variation of  
379 oxalic acid particles. Because the mass concentration of Fe (III) and oxalic acid could  
380 not be obtained through SPAMS, the diurnal variations of peak area of iron (m/z=56)  
381 and oxalic acid (m/z=-89) were used to investigate the role of iron on the net  
382 production of oxalic in the HM type particles from July 28 to August 1, 2014 (Figure  
383 8). Interestingly, the peak area of iron exhibited opposite trend with the peak area of  
384 oxalic acid from 4:00 to 11:00. As the peak area of Fe increased from 1565 to 29920  
385 from 4:00 to 7:00, the peak area of oxalic decreased from 6052 to 3487 accordingly.  
386 From 8:00 to 11:00, the peak area of Fe had a very low value of 1168, but the peak  
387 area of oxalic had a very high value of 5538. In addition, the peak area of iron  
388 exhibited a high value of 138199 at 14:00, while the peak area of oxalic acid showed a  
389 lower peak of 7687 at 14:00 and a higher peak of 11879 at 19:00 with an extremely  
390 low abundance of iron. Above opposite variation patterns of iron and oxalic acid in  
391 iron rich HM type particles during the photochemical activity period from 5:00 to  
392 19:00 strongly indicated that photolysis of iron oxalato complexes could be an  
393 efficient sink of oxalic acid.

394 The influence from traffic emission was investigated through the diurnal  
395 variations of total EC type particles and NO<sub>2</sub> (Figure S6). The EC type particles  
396 increased from 12:00 to 21:00, which had same variation as total oxalic acid, but NO<sub>2</sub>

397 followed the rush hour pattern with two peaks from 5:00 to 8:00 and from 18:00 to  
398 21:00. Traffic emission is not expected to have a large contribution to oxalic acid in  
399 this study. The wind speed was low during the whole day (Figure S6), especially  
400 between 9:00 and 18:00, which provided a stagnant environment for the increase in  
401 oxalic acid produced from photochemical process.

### 402 **3.3 Formation process of oxalic acid in winter**

403 Despite lower O<sub>3</sub> concentrations and photochemical activity in winter, oxalic  
404 acid was still prevalent in carbonaceous particles, especially BB type particles. While  
405 oxalic acid was found to be internally mixed with sulfate and nitrate both in summer  
406 and winter, the nitric acid was only observed in oxalic acid particles in winter,  
407 indicating a strongly acidic nature of oxalic acid particles in winter. Considering a  
408 possible connection of oxalic acid production with the acidic environment, the  
409 temporal concentrations of oxalic acid, sulfate and nitrate were investigated through  
410 their peak areas in the carbonaceous type oxalic acid particles including EC, OC,  
411 ECOC and BB type in Figure 9. The peaks of m/z -62[NO<sub>3</sub>]<sup>-</sup> and -97[HSO<sub>4</sub>]<sup>-</sup>  
412 represent nitrate and sulfate, respectively. Nitrate, sulfate and oxalic acid showed very  
413 similar variation patterns in winter, suggesting a close connection of the formation of  
414 oxalic acid with the existence of nitrate and sulfate. Although nitric acid was found in  
415 the oxalic acid particles, the acidity of the oxalic acid particles was not estimated  
416 since the real-time concentration of inorganic ions was not available during the  
417 sampling period in winter. Instead the relative acidity ratio (R<sub>ra</sub>), defined as the ratio  
418 of total peak areas of nitrate and sulfate to the peak area of ammonium (m/z  
419 18[NH<sub>4</sub>]<sup>+</sup>), was used (Denkenberger et al., 2007; Pratt et al., 2009). The R<sub>ra</sub> of  
420 carbonaceous type oxalic acid particles ranged from 7 to 114 with an average value of  
421 25 (Figure 9), indicating an intensely acidic environment of carbonaceous type oxalic  
422 acid particles in winter. Several studies have reported the formation of oxalic acid  
423 through the oxidation of glyoxal and related precursors in acidic aqueous phase  
424 (Carlton et al., 2006; Carlton et al., 2007; Tan et al., 2009). Although the influence of  
425 different particle acidity on the oxidation process of glyoxal still needs evaluation, the  
426 moderate acidic environment is favorable for the production of oxalic acid from the

427 oxidation of glyoxal (Herrmann, 2003;Ervens and Volkamer, 2010;Eugene et al.,  
428 2016). In this work the acidic environment of the carbonaceous type oxalic acid  
429 particles and similar variation patterns among oxalic acid, sulfate and nitrate may  
430 suggest a relationship between the degradation of organic precursors and the acidic  
431 chemical process. However, the temporal change of  $R_{ra}$  did not follow a similar trend  
432 as the peak area of oxalic acid in most particles, possibly due to the multi-step  
433 formation of oxalic acid influenced by many factors such as precursors, liquid water  
434 content and ion strength (Carlton et al., 2007;Cheng et al., 2013;Cheng et al., 2015).

435 The sharp increase of oxalic acid particles on February 8, 2015 (Figure 1) was  
436 selected as a typical episode to investigate the formation processes of oxalic acid in  
437 winter. During the episode, the 48 hr back trajectory analysis showed air masses that  
438 originated from the urban areas of Guangzhou and Foshan city (Figure S4), indicating  
439 strong influence on organic precursors from anthropogenic emissions. Oxalic acid  
440 particle types were dominated by BB (23.2%), followed by EC (22.0%) and Sec  
441 (15.1%) type (Table 2). Carbonaceous particles including EC, ECOC, OC, BB  
442 accounted for 61.6% of the total oxalic acid particles. The mass spectra of oxalic acid  
443 particles were characterized by many hydrocarbon clusters of  $27[C_2H_3]^+$ ,  $29[C_2H_5]^+$ ,  
444  $37[C_3H]^+$ ,  $43[C_2H_3O]^+$ ,  $51[C_4H_3]^+$ ,  $55[C_4H_7]^+$ ,  $63[C_3H_3]^+$ ,  $65[C_5H_5]^+$ ,  $74[C_2H_2O_3]^+$ ,  
445  $77[C_6H_5]^+$ , and carbon clusters of  $36[C_3]^+$ ,  $48[C_4]^+$ ,  $60[C_5]^+$  in positive mass spectrum,  
446 while the negative mass spectrum was characterized by elemental carbon clusters like  
447  $-24[C_2]^-$ ,  $-36[C_3]^-$ ,  $-48[C_4]^-$ , biomass burning markers of  $-59[C_2H_3O_2]^-$  and  
448  $-73[C_3H_5O_2]^-$  and secondary species including  $-42[CNO]^-$ ,  $-46[NO_2]^-$ ,  $-62[NO_3]^-$ ,  
449  $-79[PO_3]^-$ ,  $-80[SO_3]^-$ ,  $-96[SO_4]^-$  and  $-97[H_2SO_4]^-$  (Figure 10 a).

450 As the precursor of oxalic acid, glyoxal has the potential to react with sulfuric  
451 acid to produce organosulfates through acid-catalyzed nucleophilic addition according  
452 to laboratory and chamber studies(Surratt et al., 2007;Galloway et al., 2009). The  
453 negative ion of  $-155([C_2H_3O_2SO_4]^-)$  has been identified as the marker ion of  
454 organosulfates derived from glyoxal in chamber and field measurements using  
455 ATOFMS (Surratt et al., 2008;Hatch et al., 2011). The formation of organosulfates  
456 from glyoxal requires an acidic aqueous environment, which can be used as a marker

457 of acidic aqueous phase aging process of organic compounds. The temporal trend of  
458 organosulfate-containing oxalic acid particles in winter is shown in Figure S7, which  
459 exhibited a similar pattern as the total oxalic acid particles during the whole sampling  
460 period in winter. The percentage of organosulfate-containing oxalic acid particles in  
461 total oxalic acid particles ranged from 0 to 16.4% with the highest ratio observed in  
462 the episode (February 8). The linear regression between total oxalic acid particles and  
463 organosulfate-containing oxalic acid particles in the episode is exhibited in Figure 10b,  
464 and the robust correlation ( $r^2=0.81$ ) between them suggests that oxalic acid and  
465 organosulfate may share similar formation process. Based on the above discussion,  
466 the degradation of carbonaceous species associated with acidic aqueous phase  
467 chemical reactions may have an important contribution to the formation of oxalic acid  
468 during the episode in winter. Similar particle types and mass spectra of oxalic  
469 acid-containing particles during the episode and the whole sampling period in winter  
470 were observed, which suggests the acidic aqueous phase chemical processing of  
471 organic precursors as a potential source for oxalic acid.

#### 472 **4. Summary and conclusions**

473 Oxalic acid containing particles were measured by a single particle aerosol mass  
474 spectrometer (SPAMS) in the summer and winter of 2014 in Heshan, China. They  
475 accounted for 2.5% and 2.7% of the total detected ambient particles. In summer heavy  
476 metal-containing particles were the largest group of particles containing oxalic acid  
477 with a fraction of 31.3% followed by Sec type (19.2%), while in winter BB type was  
478 the dominant group with a percentage of 24.2%. More than 90% of oxalic acid  
479 particles were internally mixed with sulfate and nitrate during the whole sampling  
480 period. Only 18% of oxalic acid particles contained ammonium in summer, which  
481 increased to 71% in winter. In summer oxalic acid and O<sub>3</sub> concentration exhibited  
482 similar diurnal variations, indicating a substantial contribution of photochemical  
483 reactions to oxalic acid formation. The diurnal variations of peak area of iron and  
484 oxalic acid in HM type particles indicate a net production of oxalic acid at 15:00  
485 lower than at 19:00, likely due to a significant loss of oxalic acid through the



486 photolysis of iron oxalato complexes during the strong photochemical activity period.  
487 In winter carbonaceous type particles including EC, OC, ECOC and BB groups  
488 accounted for 59.8% of oxalic acid particles and increased to 61.6% in the episode.  
489 Nitric acid and organosulfate were found to co-exist in oxalic acid-containing  
490 particles in the winter, which suggests a close association with acidic aqueous phase  
491 reactions. Acidic aqueous phase chemical processing of organic precursors is a  
492 potential contribution for the formation of oxalic acid in winter. The current study  
493 demonstrates that SPAMS is a unique tool for understanding the mixing states of  
494 different components of ambient aerosols, which are useful for exploring the  
495 formation and evolution process of SOA.

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## 507 **References**

- 508 Aggarwal, S. G., and Kawamura, K.: Molecular distributions and stable carbon isotopic compositions  
509 of dicarboxylic acids and related compounds in aerosols from Sapporo, Japan: Implications for  
510 photochemical aging during long-range atmospheric transport, *Journal of Geophysical*  
511 *Research-Atmospheres*, 113, D14301, 10.1029/2007jd009365, 2008.
- 512 Anglada, J. M., Martins-Costa, M., Ruiz-López, M. F., and Francisco, J. S.: Spectroscopic signatures of  
513 ozone at the air–water interface and photochemistry implications, *Proceedings of the National*  
514 *Academy of Sciences*, 111, 11618-11623, 10.1073/pnas.1411727111, 2014.
- 515 Badali, K. M., Zhou, S., Aljawhary, D., Antinolo, M., Chen, W. J., Lok, A., Mungall, E., Wong, J. P. S.,  
516 Zhao, R., and Abbatt, J. P. D.: Formation of hydroxyl radicals from photolysis of secondary  
517 organic aerosol material, *Atmospheric Chemistry and Physics*, 15, 7831-7840,

518 10.5194/acp-15-7831-2015, 2015.

519 Behera, S. N., Betha, R., Liu, P., and Balasubramanian, R.: A study of diurnal variations of PM 2.5  
520 acidity and related chemical species using a new thermodynamic equilibrium model, *Science of*  
521 *the Total Environment*, 452, 286-295, 2013.

522 Bi, X., Zhang, G., Li, L., Wang, X., Li, M., Sheng, G., Fu, J., and Zhou, Z.: Mixing state of biomass  
523 burning particles by single particle aerosol mass spectrometer in the urban area of PRD, China,  
524 *Atmospheric Environment*, 45, 3447-3453, 2011.

525 Booth, A. M., Topping, D. O., McFiggans, G., and Percival, C. J.: Surface tension of mixed inorganic  
526 and dicarboxylic acid aqueous solutions at 298.15 K and their importance for cloud activation  
527 predictions, *Phys Chem Chem Phys*, 11, 8021-8028, 10.1039/b906849j, 2009.

528 Carlton, A. G., Turpin, B. J., Lim, H. J., Altieri, K. E., and Seitzinger, S.: Link between isoprene and  
529 secondary organic aerosol (SOA): Pyruvic acid oxidation yields low volatility organic acids in  
530 clouds, *Geophysical Research Letters*, 33, L06822, 10.1029/2005gl025374, 2006.

531 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S., Reff, A., Lim, H. J., and Ervens, B.:  
532 Atmospheric oxalic acid and SOA production from glyoxal: Results of aqueous photooxidation  
533 experiments, *Atmospheric Environment*, 41, 7588-7602, 10.1016/j.atmosenv.2007.05.035, 2007.

534 Cheng, C., Wang, G., Meng, J., Wang, Q., Cao, J., Li, J., and Wang, J.: Size-resolved airborne  
535 particulate oxalic and related secondary organic aerosol species in the urban atmosphere of  
536 Chengdu, China, *Atmospheric Research*, 161, 134-142, 2015.

537 Cheng, C. L., Wang, G. H., Zhou, B. H., Meng, J. J., Li, J. J., Cao, J. J., and Xiao, S.: Comparison of  
538 dicarboxylic acids and related compounds in aerosol samples collected in Xi'an, China during haze  
539 and clean periods, *Atmospheric Environment*, 81, 443-449, 10.1016/j.atmosenv.2013.09.013,  
540 2013.

541 Denkenberger, K. A., Moffet, R. C., Holecek, J. C., Rebotier, T. P., and Prather, K. A.: Real-time,  
542 single-particle measurements of oligomers in aged ambient aerosol particles, *Environmental*  
543 *Science & Technology*, 41, 5439-5446, 10.1021/es070329l, 2007.

544 Du, H., Kong, L., Cheng, T., Chen, J., Du, J., Li, L., Xia, X., Leng, C., and Huang, G.: Insights into  
545 summertime haze pollution events over Shanghai based on online water-soluble ionic composition  
546 of aerosols, *Atmospheric Environment*, 45, 5131-5137, 2011.

547 Ervens, B., Gligorovski, S., and Herrmann, H.: Temperature-dependent rate constants for hydroxyl  
548 radical reactions with organic compounds in aqueous solutions, *Phys Chem Chem Phys*, 5,  
549 1811-1824, Doi 10.1039/B300072a, 2003.

550 Ervens, B., Feingold, G., Frost, G. J., and Kreidenweis, S. M.: A modeling study of aqueous production  
551 of dicarboxylic acids: 1. Chemical pathways and speciated organic mass production, *Journal of*  
552 *Geophysical Research-Atmospheres*, 109, D15205, 10.1029/2003jd004387, 2004.

553 Ervens, B., and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry: towards a kinetic  
554 modeling framework of secondary organic aerosol formation in aqueous particles, *Atmospheric*  
555 *Chemistry and Physics*, 10, 8219-8244, DOI 10.5194/acp-10-8219-2010, 2010.

556 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in cloud droplets and  
557 aqueous particles (aqSOA): a review of laboratory, field and model studies, *Atmospheric*  
558 *Chemistry and Physics*, 11, 11069-11102, 10.5194/acp-11-11069-2011, 2011.

559 Ervens, B., Sorooshian, A., Lim, Y. B., and Turpin, B. J.: Key parameters controlling OH-initiated  
560 formation of secondary organic aerosol in the aqueous phase (aqSOA), *Journal of Geophysical*  
561 *Research-Atmospheres*, 119, 3997-4016, 10.1002/2013JD021021, 2014.

562 Ervens, B.: Modeling the Processing of Aerosol and Trace Gases in Clouds and Fogs, *Chem Rev*, 115,  
563 4157-4198, 10.1021/cr5005887, 2015.

564 Eugene, A. J., Xia, S. S., and Guzman, M. I.: Aqueous Photochemistry of Glyoxylic Acid, *Journal of*  
565 *Physical Chemistry A*, 120, 3817-3826, 10.1021/acs.jpca.6b00225, 2016.

566 Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient thermodynamic equilibrium  
567 model for  $K^+$ - $Ca^{2+}$ - $Mg^{2+}$ - $NH_4^+$ - $Na^+$ - $SO_4^{2-}$ - $NO_3^-$ - $Cl^-$ - $H_2O$  aerosols, *Atmospheric*  
568 *Chemistry and Physics*, 7, 4639-4659, 2007.

569 Galloway, M. M., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Flagan, R. C., Seinfeld, J. H., and  
570 Keutsch, F. N.: Glyoxal uptake on ammonium sulphate seed aerosol: reaction products and  
571 reversibility of uptake under dark and irradiated conditions, *Atmospheric Chemistry and Physics*,  
572 9, 3331-3345, 10.5194/acp-9-3331-2009, 2009.

573 Gligorovski, S., Strekowski, R., Barbati, S., and Vione, D.: Environmental Implications of Hydroxyl  
574 Radicals ( $\bullet OH$ ), *Chem Rev*, 115, 13051-13092, 10.1021/cr500310b, 2015.

575 Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in the earth's  
576 atmosphere, *Environmental Science & Technology*, 41, 1514-1521, Doi 10.1021/Es072476p,  
577 2007.

578 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M., Dommen, J.,  
579 Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F., Herrmann, H., Hoffmann, T.,  
580 Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L., Kiendler-Scharr, A., Maenhaut, W., McFiggans,  
581 G., Mentel, T. F., Monod, A., Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and  
582 Wildt, J.: The formation, properties and impact of secondary organic aerosol: current and  
583 emerging issues, *Atmospheric Chemistry and Physics*, 9, 5155-5236, 10.5194/acp-9-5155-2009,  
584 2009.

585 Hatch, L. E., Creamean, J. M., Ault, A. P., Surratt, J. D., Chan, M. N., Seinfeld, J. H., Edgerton, E. S.,  
586 Su, Y., and Prather, K. A.: Measurements of isoprene-derived organosulfates in ambient aerosols  
587 by aerosol time-of-flight mass spectrometry-Part 1: Single particle atmospheric observations in  
588 Atlanta, *Environmental science & technology*, 45, 5105-5111, 2011.

589 Herrmann, H.: Kinetics of aqueous phase reactions relevant for atmospheric chemistry, *Chem Rev*, 103,  
590 4691-4716, Doi 10.1021/Cr020658q, 2003.

591 Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M., and Otto, T.: Tropospheric  
592 Aqueous-Phase Chemistry: Kinetics, Mechanisms, and Its Coupling to a Changing Gas Phase,  
593 *Chem Rev*, 115, 4259-4334, 10.1021/cr500447k, 2015.

594 Ho, K. F., Cao, J. J., Lee, S. C., Kawamura, K., Zhang, R. J., Chow, J. C., and Watson, J. G.:  
595 Dicarboxylic acids, ketocarboxylic acids, and dicarbonyls in the urban atmosphere of China,  
596 *Journal of Geophysical Research-Atmospheres*, 112, D22S27, 10.1029/2006jd008011, 2007.

597 Ho, K. F., Lee, S. C., Ho, S. S. H., Kawamura, K., Tachibana, E., Cheng, Y., and Zhu, T.: Dicarboxylic  
598 acids, ketocarboxylic acids, alpha-dicarbonyls, fatty acids, and benzoic acid in urban aerosols  
599 collected during the 2006 Campaign of Air Quality Research in Beijing (CAREBeijing-2006),  
600 *Journal of Geophysical Research-Atmospheres*, 115, D19312, 10.1029/2009jd013304, 2010.

601 Ho, K. F., Ho, S. S. H., Lee, S. C., Kawamura, K., Zou, S. C., Cao, J. J., and Xu, H. M.: Summer and  
602 winter variations of dicarboxylic acids, fatty acids and benzoic acid in PM(2.5) in Pearl Delta  
603 River Region, China, *Atmospheric Chemistry and Physics*, 11, 2197-2208,  
604 10.5194/acp-11-2197-2011, 2011.

605 Huang, X.-F., and Yu, J. Z.: Is vehicle exhaust a significant primary source of oxalic acid in ambient

606 aerosols?, *Geophysical Research Letters*, 34, L02808, 10.1029/2006gl028457, 2007.

607 Jacob, D. J.: Chemistry of OH in remote clouds and its role in the production of formic acid and  
608 peroxymonosulfate, *Journal of Geophysical Research: Atmospheres*, 91, 9807-9826,  
609 10.1029/JD091iD09p09807, 1986.

610 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q., Kroll, J. H., DeCarlo,  
611 P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K. S., Ulbrich, I. M., Grieshop, A. P.,  
612 Robinson, A. L., Duplissy, J., Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L.,  
613 Tian, J., Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M.,  
614 Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J. A., Onasch, T. B.,  
615 Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S.,  
616 Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T.,  
617 Hatakeyama, S., Shimojo, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D.,  
618 Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A. M.,  
619 Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic Aerosols in the  
620 Atmosphere, *Science*, 326, 1525-1529, 10.1126/science.1180353, 2009.

621 Kawamura, K., and Ikushima, K.: seasonal-changes in the distribution of dicarboxylic-acids in the  
622 urban atmosphere, *Environmental Science & Technology*, 27, 2227-2235, 10.1021/es00047a033,  
623 1993.

624 Kawamura, K., Kobayashi, M., Tsubonuma, N., Mochida, M., Watanabe, T., and Lee, M.: Organic and  
625 inorganic compositions of marine aerosols from East Asia: Seasonal variations of water-soluble  
626 dicarboxylic acids, major ions, total carbon and nitrogen, and stable C and N isotopic composition,  
627 *Geochemical Investigations in Earth and Space Science: A Tribute to Issac R. Kaplan*, 9, edited by:  
628 Hill, R. J. L. J. A. Z. B. M. J. C. G., and Eganhouse, R. G. M. P. K., 243-265 pp., 2004.

629 Kawamura, K., and Yasui, O.: Diurnal changes in the distribution of dicarboxylic acids, ketocarboxylic  
630 acids and dicarbonyls in the urban Tokyo atmosphere, *Atmospheric Environment*, 39, 1945-1960,  
631 10.1016/j.atmosenv.2004.12.014, 2005.

632 Kawamura, K., Kasukabe, H., and Barrie, L. A.: Secondary formation of water-soluble organic acids  
633 and alpha-dicarbonyls and their contributions to total carbon and water-soluble organic carbon:  
634 Photochemical aging of organic aerosols in the Arctic spring, *Journal of Geophysical  
635 Research-Atmospheres*, 115, D21306, 10.1029/2010jd014299, 2010.

636 Kawamura, K., Tachibana, E., Okuzawa, K., Aggarwal, S. G., Kanaya, Y., and Wang, Z. F.: High  
637 abundances of water-soluble dicarboxylic acids, ketocarboxylic acids and alpha-dicarbonyls in the  
638 mountaintop aerosols over the North China Plain during wheat burning season, *Atmospheric  
639 Chemistry and Physics*, 13, 8285-8302, 10.5194/acp-13-8285-2013, 2013.

640 Khezri, B., Mo, H., Yan, Z., Chong, S.-L., Heng, A. K., and Webster, R. D.: Simultaneous online  
641 monitoring of inorganic compounds in aerosols and gases in an industrialized area, *Atmospheric  
642 Environment*, 80, 352-360, 2013.

643 Kundu, S., Kawamura, K., Andreae, T. W., Hoffer, A., and Andreae, M. O.: Molecular distributions of  
644 dicarboxylic acids, ketocarboxylic acids and alpha-dicarbonyls in biomass burning aerosols:  
645 implications for photochemical production and degradation in smoke layers, *Atmospheric  
646 Chemistry and Physics*, 10, 2209-2225, 10.5194/acp-10-2209-2010, 2010.

647 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H., Bi, X. H., Cheng,  
648 P., and Zhou, Z.: Real time bipolar time-of-flight mass spectrometer for analyzing single aerosol  
649 particles, *Int J Mass Spectrom*, 303, 118-124, 10.1016/j.ijms.2011.01.017, 2011.

650 Li, L., Li, M., Huang, Z., Gao, W., Nian, H., Fu, Z., Gao, J., Chai, F., and Zhou, Z.: Ambient particle  
651 characterization by single particle aerosol mass spectrometry in an urban area of Beijing,  
652 *Atmospheric Environment*, 94, 323-331, 2014.

653 Lim, Y. B., Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Aqueous chemistry and its role in  
654 secondary organic aerosol (SOA) formation, *Atmospheric Chemistry and Physics*, 10,  
655 10521-10539, 10.5194/acp-10-10521-2010, 2010.

656 Ma, Q. X., He, H., and Liu, C.: Hygroscopic properties of oxalic acid and atmospherically relevant  
657 oxalates, *Atmospheric Environment*, 69, 281-288, 10.1016/j.atmosenv.2012.12.011, 2013.

658 McNeill, V. F.: Aqueous organic chemistry in the atmosphere: Sources and chemical processing of  
659 organic aerosols, *Environmental science & technology*, 49, 1237-1244, 2015.

660 Meng, J. J., Wang, G. H., Li, J. J., Cheng, C. L., and Cao, J. J.: Atmospheric oxalic acid and related  
661 secondary organic aerosols in Qinghai Lake, a continental background site in Tibet Plateau,  
662 *Atmospheric Environment*, 79, 582-589, 10.1016/j.atmosenv.2013.07.024, 2013.

663 Meng, J. J., Wang, G. H., Li, J. J., Cheng, C. L., Ren, Y. Q., Huang, Y., Cheng, Y. T., Cao, J. J., and  
664 Zhang, T.: Seasonal characteristics of oxalic acid and related SOA in the free troposphere of Mt.  
665 Hua, central China: Implications for sources and formation mechanisms, *Science Of The Total  
666 Environment*, 493, 1088-1097, 10.1016/j.scitotenv.2014.04.086, 2014.

667 Miyazaki, Y., Aggarwal, S. G., Singh, K., Gupta, P. K., and Kawamura, K.: Dicarboxylic acids and  
668 water-soluble organic carbon in aerosols in New Delhi, India, in winter: Characteristics and  
669 formation processes, *Journal of Geophysical Research-Atmospheres*, 114, D19206,  
670 10.1029/2009jd011790, 2009.

671 Myriokefalitakis, S., Tsigaridis, K., Mihalopoulos, N., Sciare, J., Nenes, A., Kawamura, K., Segers, A.,  
672 and Kanakidou, M.: In-cloud oxalate formation in the global troposphere: a 3-D modeling study,  
673 *Atmospheric Chemistry and Physics*, 11, 5761-5782, 10.5194/acp-11-5761-2011, 2011.

674 Nenes, A., Pandis, S. N., and Pilinis, C.: ISORROPIA: A new thermodynamic equilibrium model for  
675 multiphase multicomponent inorganic aerosols, *Aquatic geochemistry*, 4, 123-152, 1998.

676 Nenes, A., Pandis, S. N., and Pilinis, C.: Continued development and testing of a new thermodynamic  
677 aerosol module for urban and regional air quality models, *Atmospheric Environment*, 33,  
678 1553-1560, 1999.

679 Noble, C. A., and Prather, K. A.: Real-time measurement of correlated size and composition profiles of  
680 individual atmospheric aerosol particles, *Environmental science & technology*, 30, 2667-2680,  
681 1996.

682 Novakov, T., and Penner, J. E.: Large Contribution of Organic Aerosols to Cloud-Condensation-Nuclei  
683 Concentrations, *Nature*, 365, 823-826, Doi 10.1038/365823a0, 1993.

684 Pavuluri, C. M., Kawamura, K., and Swaminathan, T.: Water-soluble organic carbon, dicarboxylic acids,  
685 ketoacids, and alpha-dicarbonyls in the tropical Indian aerosols, *Journal of Geophysical  
686 Research-Atmospheres*, 115, D11302, 10.1029/2009JD012661, 2010.

687 Poschl, U., and Shiraiwa, M.: Multiphase Chemistry at the Atmosphere-Biosphere Interface Influencing  
688 Climate and Public Health in the Anthropocene, *Chem Rev*, 115, 4440-4475, 10.1021/cr500487s,  
689 2015.

690 Prather, K. A., Nordmeyer, T., and Salt, K.: Real-time characterization of individual aerosol particles  
691 using time-of-flight mass spectrometry, *Anal. Chem.*, 66, 1403-1407, 1994.

692 Pratt, K. A., Hatch, L. E., and Prather, K. A.: Seasonal Volatility Dependence of Ambient Particle Phase  
693 Amines, *Environmental Science & Technology*, 43, 5276-5281, 10.1021/es803189n, 2009.

694 Prenni, A. J., De Mott, P. J., and Kreidenweis, S. M.: Water uptake of internally mixed particles  
695 containing ammonium sulfate and dicarboxylic acids, *Atmospheric Environment*, 37, 4243-4251,  
696 10.1016/s1352-2310(03)00559-4, 2003.

697 Silva, P. J., and Prather, K. A.: Interpretation of mass spectra from organic compounds in aerosol  
698 time-of-flight mass spectrometry, *Anal. Chem.*, 72, 3553-3562, 2000.

699 Sorooshian, A., Varutbangkul, V., Brechtel, F. J., Ervens, B., Feingold, G., Bahreini, R., Murphy, S. M.,  
700 Holloway, J. S., Atlas, E. L., Buzorius, G., Jonsson, H., Flagan, R. C., and Seinfeld, J. H.: Oxalic  
701 acid in clear and cloudy atmospheres: Analysis of data from International Consortium for  
702 Atmospheric Research on Transport and Transformation 2004, *Journal of Geophysical  
703 Research-Atmospheres*, 111, 10.1029/2005jd006880, 2006.

704 Sorooshian, A., Lu, M.-L., Brechtel, F. J., Jonsson, H., Feingold, G., Flagan, R. C., and Seinfeld, J. H.:  
705 On the source of organic acid aerosol layers above clouds, *Environmental Science & Technology*,  
706 41, 4647-4654, 10.1021/es0630442, 2007a.

707 Sorooshian, A., Ng, N. L., Chan, A. W. H., Feingold, G., Flagan, R. C., and Seinfeld, J. H.: Particulate  
708 organic acids and overall water-soluble aerosol composition measurements from the 2006 Gulf of  
709 Mexico Atmospheric Composition and Climate Study (GoMACCS), *Journal of Geophysical  
710 Research-Atmospheres*, 112, 10.1029/2007jd008537, 2007b.

711 Sorooshian, A., Wang, Z., Coggon, M. M., Jonsson, H. H., and Ervens, B.: Observations of Sharp  
712 Oxalate Reductions in Stratocumulus Clouds at Variable Altitudes: Organic Acid and Metal  
713 Measurements During the 2011 E-PEACE Campaign, *Environmental Science & Technology*, 47,  
714 7747-7756, 10.1021/es4012383, 2013.

715 Stone, E. A., Hedman, C. J., Zhou, J. B., Mieritz, M., and Schauer, J. J.: Insights into the nature of  
716 secondary organic aerosol in Mexico City during the MILAGRO experiment 2006, *Atmospheric  
717 Environment*, 44, 312-319, 10.1016/j.atmosenv.2009.10.036, 2010.

718 Sullivan, R. C., and Prather, K. A.: Investigations of the diurnal cycle and mixing state of oxalic acid in  
719 individual particles in Asian aerosol outflow, *Environmental Science & Technology*, 41,  
720 8062-8069, 10.1021/es071134g, 2007a.

721 Sullivan, R. C., and Prather, K. A.: Investigations of the diurnal cycle and mixing state of oxalic acid in  
722 individual particles in Asian aerosol outflow, *Environmental Science Technology*, 41, 8062-8069,  
723 2007b.

724 Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A., Ng, N. L.,  
725 Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C., and Seinfeld, J. H.: Evidence for  
726 organosulfates in secondary organic aerosol, *Environmental Science & Technology*, 41, 517-527,  
727 10.1021/Es062081q, 2007.

728 Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeylen, R., Shahgholi, M., Kleindienst, T. E.,  
729 Edney, E. O., Offenberg, J. H., Lewandowski, M., Jaoui, M., Maenhaut, W., Claeys, M., Flagan, R.  
730 C., and Seinfeld, J. H.: Organosulfate formation in biogenic secondary organic aerosol, *Journal of  
731 Physical Chemistry A*, 112, 8345-8378, Doi 10.1021/Jp802310p, 2008.

732 Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Effects of Precursor Concentration and Acidic  
733 Sulfate in Aqueous Glyoxal-OH Radical Oxidation and Implications for Secondary Organic  
734 Aerosol, *Environmental Science & Technology*, 43, 8105-8112, 10.1021/Es901742f, 2009.

735 ten Brink, H., Otjes, R., Jongejan, P., and Slanina, S.: An instrument for semi-continuous monitoring of  
736 the size-distribution of nitrate, ammonium, sulphate and chloride in aerosol, *Atmospheric  
737 Environment*, 41, 2768-2779, 2007.

738 Tong, H. J., Arangio, A. M., Lakey, P. S. J., Berkemeier, T., Liu, F. B., Kampf, C. J., Brune, W. H.,  
739 Poschl, U., and Shiraiwa, M.: Hydroxyl radicals from secondary organic aerosol decomposition in  
740 water, *Atmospheric Chemistry and Physics*, 16, 1761-1771, 10.5194/acp-16-1761-2016, 2016.

741 van Pinxteren, D., Neususs, C., and Herrmann, H.: On the abundance and source contributions of  
742 dicarboxylic acids in size-resolved aerosol particles at continental sites in central Europe,  
743 *Atmospheric Chemistry and Physics*, 14, 3913-3928, 10.5194/acp-14-3913-2014, 2014.

744 Wang, G., Cheng, C., Meng, J., Huang, Y., Li, J., and Ren, Y.: Field observation on secondary organic  
745 aerosols during Asian dust storm periods: Formation mechanism of oxalic acid and related  
746 compounds on dust surface, *Atmospheric Environment*, 113, 169-176, 2015.

747 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y., Peng, J., Guo, S.,  
748 Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J., Cao, J., An, Z., Zhou, W., Li, G.,  
749 Wang, J., Tian, P., Marrero-Ortiz, W., Secret, J., Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M.,  
750 Wang, W., Huang, Y., Wang, Y., Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F.,  
751 Rosenfeld, D., Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate formation  
752 from London Fog to Chinese haze, *Proceedings of the National Academy of Sciences*, 113,  
753 13630-13635, 10.1073/pnas.1616540113, 2016.

754 Wang, G. H., Kawamura, K., Cheng, C. L., Li, J. J., Cao, J. J., Zhang, R. J., Zhang, T., Liu, S. X., and  
755 Zhao, Z. Z.: Molecular Distribution and Stable Carbon Isotopic Composition of Dicarboxylic  
756 Acids, Ketocarboxylic Acids, and alpha-Dicarbonyls in Size-Resolved Atmospheric Particles  
757 From Xi'an City, China, *Environmental Science & Technology*, 46, 4783-4791,  
758 10.1021/es204322c, 2012.

759 Wang, J. Y., Wang, G. H., Gao, J., Wang, H., Ren, Y. Q., Li, J. J., Wu, C., Zhang, L., Wang, S. L., and  
760 Chai, F. H.: Concentrations and stable carbon isotope compositions of oxalic acid and related SOA  
761 in Beijing before, during and after the 2014 APEC, *Atmospheric Chemistry and Physics* 17,  
762 981-992, 2017.

763 Weller, C., Tilgner, A., Brauer, P., and Herrmann, H.: Modeling the Impact of Iron-Carboxylate  
764 Photochemistry on Radical Budget and Carboxylate Degradation in Cloud Droplets and Particles,  
765 *Environmental Science & Technology*, 48, 5652-5659, 10.1021/es4056643, 2014.

766 Wonaschuetz, A., Sorooshian, A., Ervens, B., Chuang, P. Y., Feingold, G., Murphy, S. M., de Gouw, J.,  
767 Warneke, C., and Jonsson, H. H.: Aerosol and gas re-distribution by shallow cumulus clouds: An  
768 investigation using airborne measurements, *Journal of Geophysical Research-Atmospheres*, 117,  
769 10.1029/2012jd018089, 2012.

770 Xue, J., Lau, A. K. H., and Yu, J. Z.: A study of acidity on PM2.5 in Hong Kong using online ionic  
771 chemical composition measurements, *Atmospheric Environment*, 45, 7081-7088,  
772 10.1016/j.atmosenv.2011.09.040, 2011.

773 Yang, F., Chen, H., Wang, X., Yang, X., Du, J., and Chen, J.: Single particle mass spectrometry of  
774 oxalic acid in ambient aerosols in Shanghai: Mixing state and formation mechanism, *Atmospheric  
775 Environment*, 43, 3876-3882, 2009.

776 Yao, X. H., Fang, M., and Chan, C. K.: Size distributions and formation of dicarboxylic acids in  
777 atmospheric particles, *Atmospheric Environment*, 36, 2099-2107, 2002.

778 Yao, X. H., Lau, A. P. S., Fang, M., Chan, C. K., and Hu, M.: Size distributions and formation of ionic  
779 species in atmospheric particulate pollutants in Beijing, China: 2 - dicarboxylic acids,  
780 *Atmospheric Environment*, 37, 3001-3007, 10.1016/s1352-2310(03)00256-5, 2003.

781 Yu, J. Z., Huang, X. F., Xu, J. H., and Hu, M.: When aerosol sulfate goes up, so does oxalate:

782 Implication for the formation mechanisms of oxalate, *Environmental Science & Technology*, 39,  
783 128-133, 10.1021/Es049559f, 2005.

784 Yu, L., Smith, J., Laskin, A., Anastasio, C., Laskin, J., and Zhang, Q.: Chemical characterization of  
785 SOA formed from aqueous-phase reactions of phenols with the triplet excited state of carbonyl  
786 and hydroxyl radical, *Atmospheric Chemistry and Physics*, 14, 13801-13816,  
787 10.5194/acp-14-13801-2014, 2014.

788 Zauscher, M. D., Wang, Y., Moore, M. J. K., Gaston, C. J., and Prather, K. A.: Air Quality Impact and  
789 Physicochemical Aging of Biomass Burning Aerosols during the 2007 San Diego Wildfires,  
790 *Environmental Science & Technology*, 47, 7633-7643, 10.1021/es4004137, 2013.

791 Zhang, G., Bi, X., Li, L., Chan, L. Y., Li, M., Wang, X., Sheng, G., Fu, J., and Zhou, Z.: Mixing state of  
792 individual submicron carbon-containing particles during spring and fall seasons in urban  
793 Guangzhou, China: a case study, *Atmospheric Chemistry and Physics*, 13, 4723-4735, 2013.

794 Zhang, G., Bi, X., He, J., Chen, D., Chan, L. Y., Xie, G., Wang, X., Sheng, G., Fu, J., and Zhou, Z.:  
795 Variation of secondary coatings associated with elemental carbon by single particle analysis,  
796 *Atmospheric Environment*, 92, 162-170, 2014.

797 Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M., and Wang, Y.:  
798 Formation of urban fine particulate matter, *Chem Rev*, 115, 3803-3855, 2015.

799 Zhou, Y., Huang, X. H., Bian, Q., Griffith, S. M., Louie, P. K., and Yu, J. Z.: Sources and atmospheric  
800 processes impacting oxalate at a suburban coastal site in Hong Kong: Insights inferred from 1 year  
801 hourly measurements, *Journal of Geophysical Research: Atmospheres*, 120, 9772-9788, 2015.

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826 **Tables and Figures**

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828 **Table list:**

829

830 Table 1. Summary of major groups of oxalic acid-containing particles in summer and  
831 winter in PRD, China.

832

833 Table 2. The abundance of major particle types in total oxalic acid-containing  
834 particles during the episode in winter (2/8/2015).

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836 **Figure caption:**

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838 Figure 1. Temporal variations of total detected particles and oxalic acid containing  
839 particles during whole sampling periods in Heshan, China: (a) hourly variations of  
840 PM<sub>2.5</sub> mass concentration, total detected particle counts, oxalic acid containing  
841 particles, ratio of oxalic acid-containing/total particles and major types of oxalic acid  
842 containing particles; (b) variation patterns of relative abundance of major types of  
843 oxalic acid containing particles.

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845 Figure 2. The averaged positive and negative ion mass spectra of oxalic acid  
846 containing particles is investigated in summer and winter: (a) summer positive, (b)  
847 summer negative, (c) winter positive, (d) winter negative. The color bars represent  
848 each peak area corresponding to specific fraction in individual particles.

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850 Figure 3. (a) Mixing state of oxalic acid with sulfate, nitrate and ammonium in oxalic  
851 acid-containing particles; (b) Linear correlation between NH<sub>4</sub><sup>+</sup>-containing oxalic acid  
852 particles and the total oxalic acid particles in summer; (c) Linear correlation between  
853 NH<sub>4</sub><sup>+</sup>-containing oxalic acid particles and the total oxalic acid particles in winter.  
854 Abbreviations: C<sub>2</sub>-NH<sub>4</sub><sup>+</sup> represents the NH<sub>4</sub><sup>+</sup>-containing oxalic acid particles, and  
855 same expressions for C<sub>2</sub>-SO<sub>4</sub><sup>2-</sup> and C<sub>2</sub>-NO<sub>3</sub><sup>-</sup>.

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857 Figure 4. Unscaled size-resolved number distributions of major types of oxalic acid  
858 particles in summer and winter.

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860 Figure 5. Temporal variations of O<sub>3</sub> concentrations, oxalic acid particles, malonic acid  
861 particles and heavy metal type of oxalic acid particles during the entire sampling  
862 period in Heshan, China.

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864 Figure 6. The averaged digitized positive and negative ion mass spectra of heavy  
865 metal type of oxalic acid-containing particles in summer.

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867 Figure 7. The diurnal variations of O<sub>3</sub> concentration, oxalic acid particles, HM group  
868 particles and in-situ pH ( $pH_{is}$ ) from July 28 to August 1 in 2014.

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870 Figure 8. The diurnal variations of peak area of iron ( $m/z=56$ ) and oxalic acid  
871 ( $m/z=-89$ ) in the HM type oxalic acid particles from July 28 to August 1, 2014.

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873 Figure 9. The temporal variations of peak area of nitrate, sulfate and oxalic acid, and  
874 the relative acidity ratio ( $R_{ra}$ ) in carbonaceous type oxalic acid particles in winter.

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876 Figure 10. The comprehensive study of oxalic acid particles increase on Feb 8, 2015:  
877 (a) The digitized positive and negative ion mass spectrum of oxalic acid particles  
878 during the episode; (b) Linear regression between total oxalic acid particles and  
879 organosulfate-containing oxalic acid particles ( $m/z -155$ ).

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Table 1. Summary of major groups of oxalic acid-containing particles in summer and winter in PRD, China.

Particle type	Summer(7/18-8/1, 2014)		Winter(1/27-2/8, 2015)	
	Count	Percentage, %	Count	Percentage, %
EC	1473	11.2	3161	15.4
ECOC	41	0.3	2233	10.9
OC	473	3.6	1922	9.4
BB	1702	13.0	4953	24.2
HM	4104	31.3	3124	15.2
Sec	2511	19.2	2192	10.7
NaK	303	2.3	17	0.1
Dust	1139	8.7	1888	9.2

Abbreviations of major particle types: elemental carbon (EC), elemental and organic carbon (ECOC), organic carbon (OC), biomass burning (BB), heavy metal (HM), secondary (Sec), sodium and potassium (NaK) and dust (Dust).

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Table 2. The abundance of major particle types in total oxalic acid-containing particles during the episode in winter (2/8/2015).

	EC	ECOC	OC	BB	Sec	HM	Dust	other
Count	1250	604	326	1320	856	377	814	132
Percentage, %	22.0	10.6	5.7	23.2	15.1	6.6	14.3	2.3

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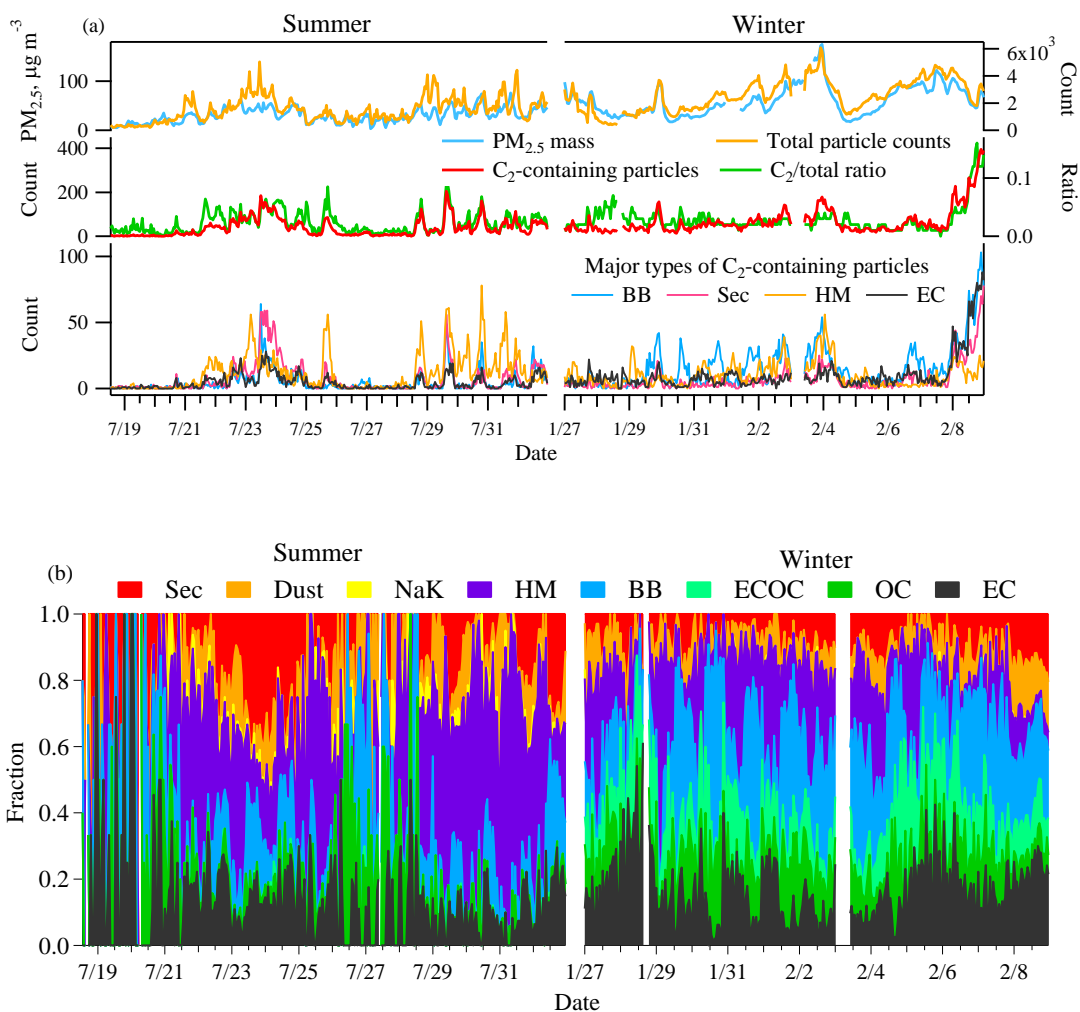
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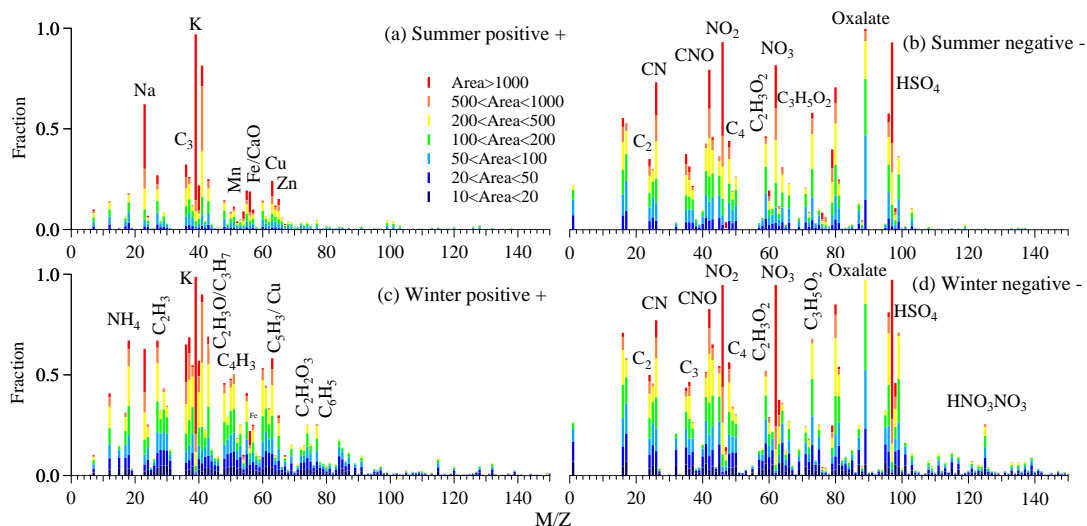


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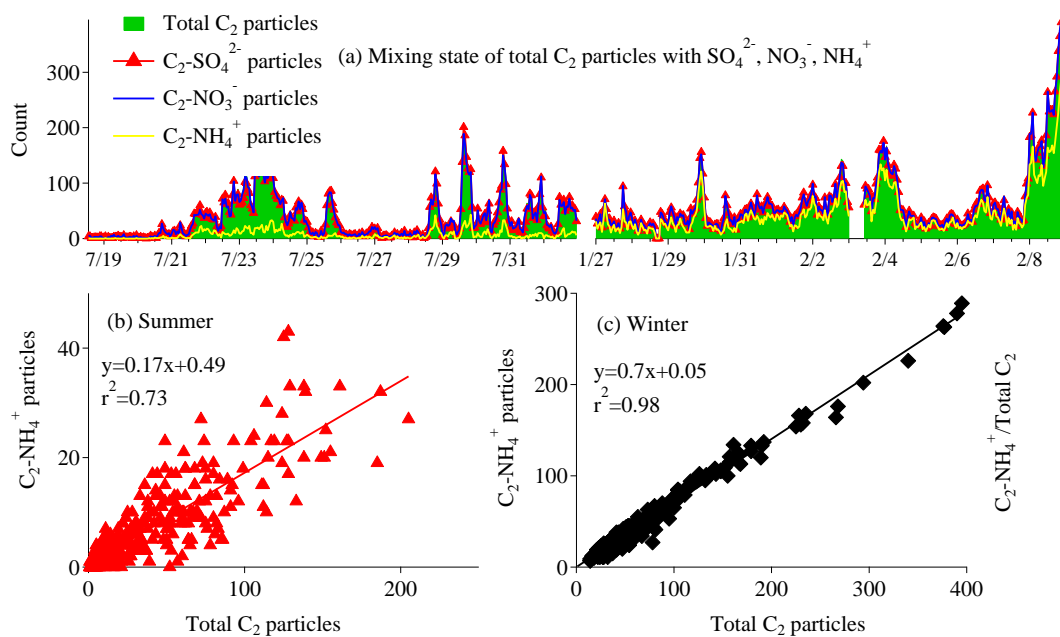
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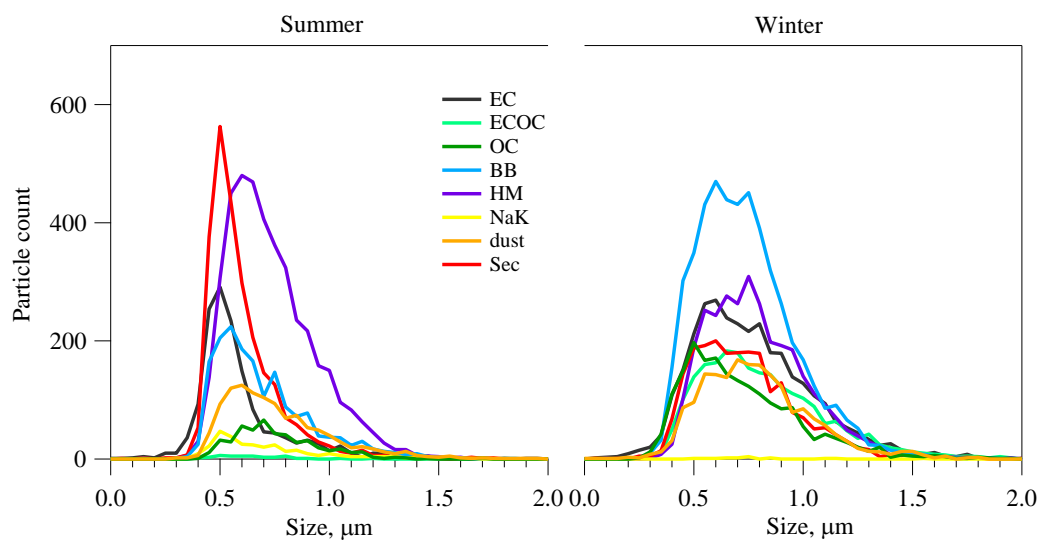
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981 Figure 4. Unscaled size-resolved number distributions of major types of oxalic acid  
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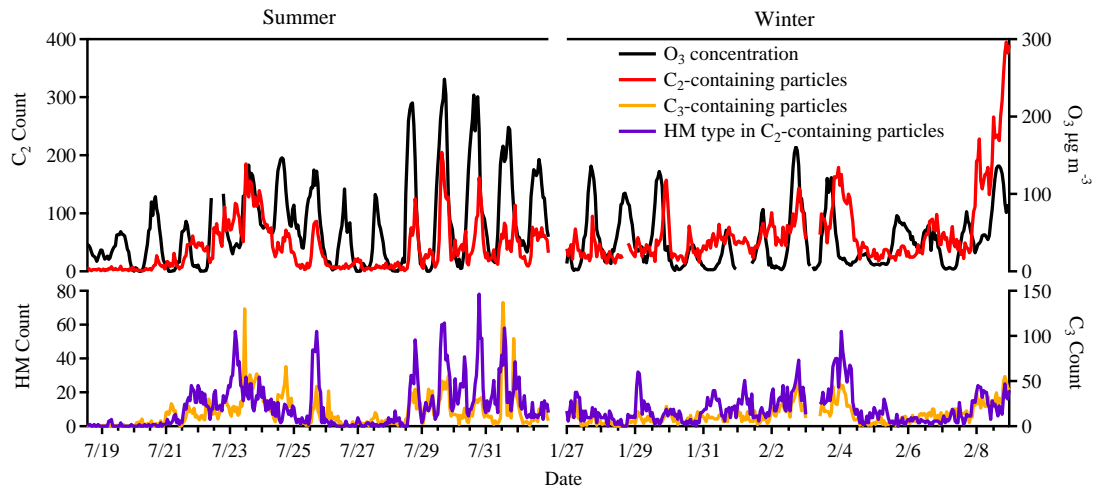
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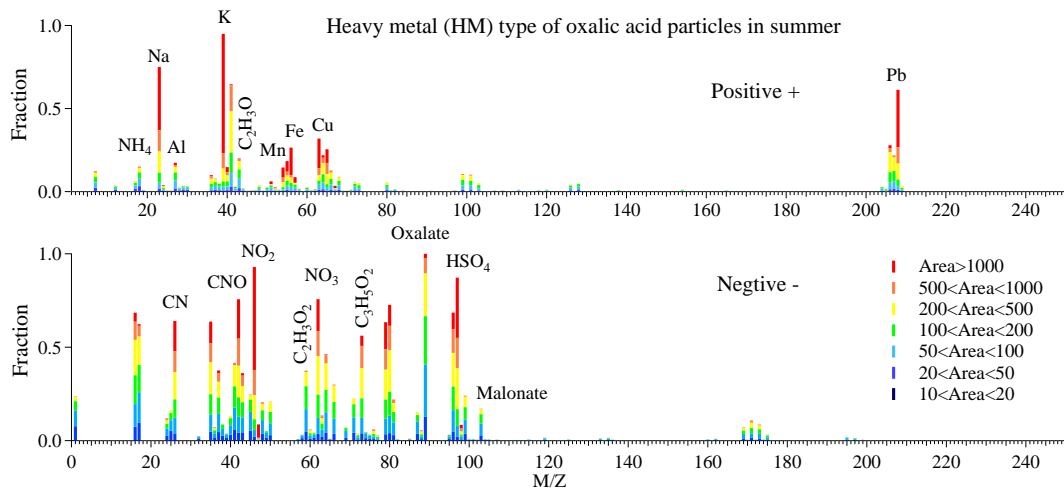
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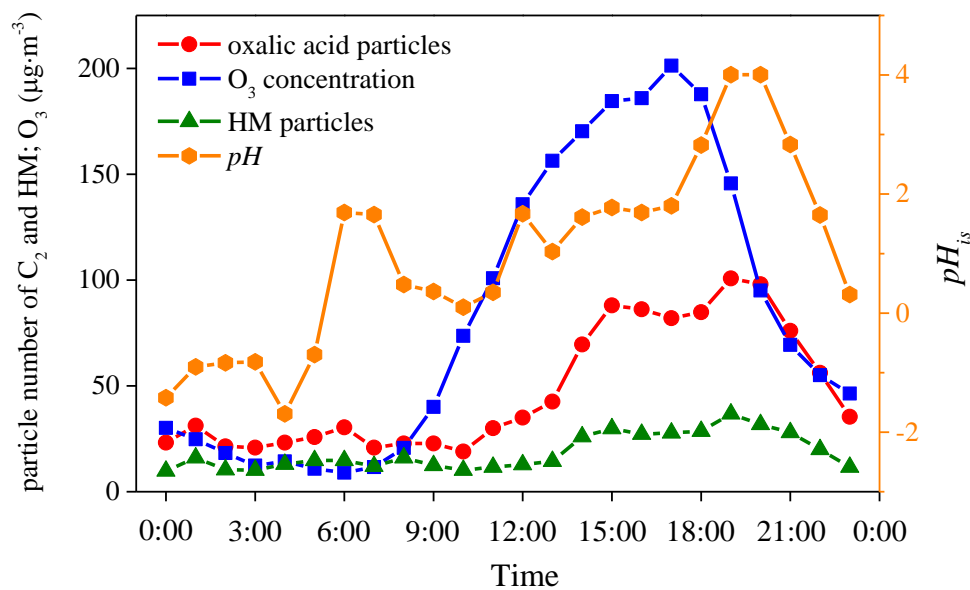
Figure 5. Temporal variations of O<sub>3</sub> concentrations, oxalic acid particles, malonic acid particles and heavy metal type of oxalic acid particles during the entire sampling period in Heshan, China.



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Figure 6. The averaged digitized positive and negative ion mass spectra of heavy metal type of oxalic acid-containing particles in summer.





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1022 Figure 7. The diurnal variations of O<sub>3</sub> concentration, oxalic acid particles, HM group  
 1023 particles and in-situ pH ( $pH_{is}$ ) from July 28 to August 1 in 2014.

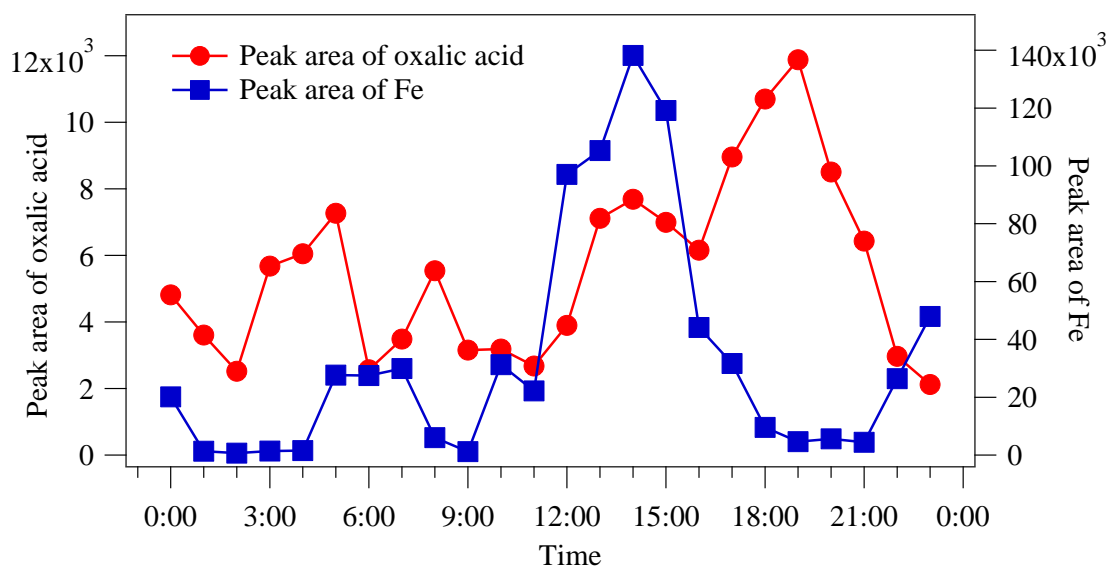
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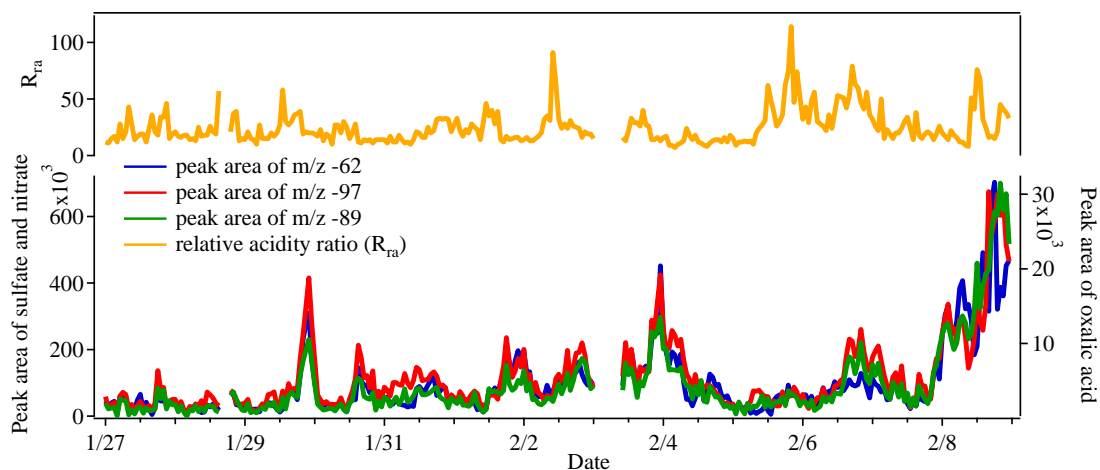
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1030 Figure 8. The diurnal variations of peak area of iron ( $m/z=56$ ) and oxalic acid  
 1031 ( $m/z=-89$ ) in the HM type oxalic acid particles from July 28 to August 1, 2014.

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1035 Figure 9. The temporal variations of peak area of nitrate, sulfate and oxalic acid, and  
 1036 the relative acidity ratio ( $R_{ra}$ ) in carbonaceous type oxalic acid particles in winter.

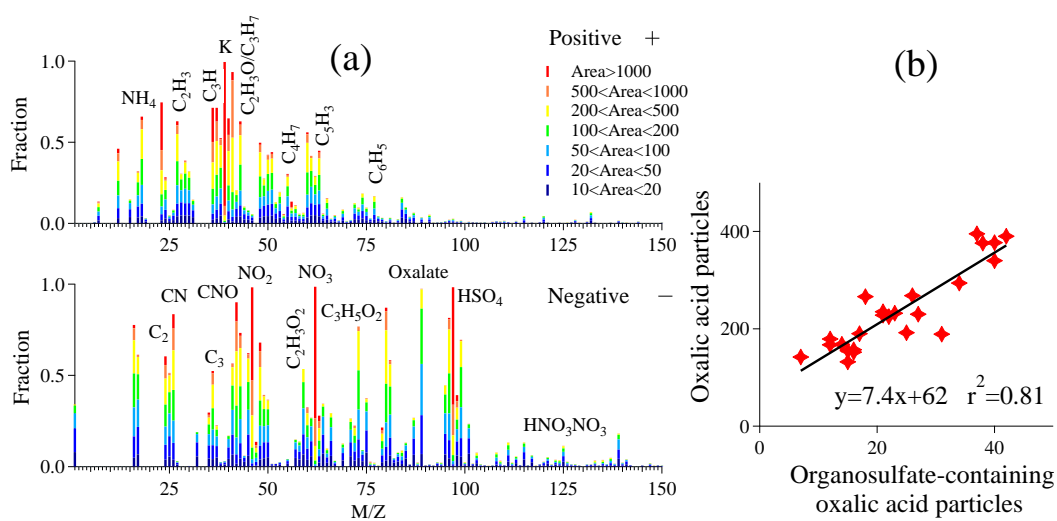
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1043 Figure 10. The comprehensive study of oxalic acid particles increase on Feb 8, 2015:

1044 (a) The digitized positive and negative ion mass spectrum of oxalic acid particles

1045 during the episode; (b) Linear regression between total oxalic acid particles and

1046 organosulfate-containing oxalic acid particles ( $m/z$  -155).

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