

1 **Mixing state of oxalic acid containing particles in the rural area of Pearl**
2 **River Delta, China: implications for the formation mechanism of oxalic**
3 **acid**

4
5 Chunlei Cheng^{1,2}, Mei Li^{1,2*}, Chak K. Chan³, Haijie Tong⁴, Changhong Chen⁵,
6 Duohong Chen⁶, Dui Wu^{1,2}, Lei Li^{1,2}, Cheng Wu^{1,2}, Peng Cheng^{1,2}, Wei Gao^{1,2},
7 Zhengxu Huang^{1,2}, Xue Li^{1,2}, Zhijuan Zhang^{1,2}, Zhong Fu⁷, Yanru Bi⁷, Zhen Zhou^{1,2*}

8
9
10 ¹Institute of Mass Spectrometer and Atmospheric Environment, Jinan University,
11 Guangzhou 510632, China

12 ²Guangdong Provincial Engineering Research Center for on-line source apportionmen
13 t system of air pollution, Guangzhou 510632, China

14 ³School of Energy and Environment, City University of Hong Kong, Hong Kong,
15 China

16 ⁴Max Planck Institute for Chemistry, Multiphase Chemistry Department,
17 Hahn-Meitner-Weg 1, 55128 Mainz, Germany

18 ⁵State of Environmental Protection Key Laboratory of the formation and prevention of
19 urban air pollution complex, Shanghai Academy of Environmental Sciences, Shanghai
20 200233, China

21 ⁶State Environmental Protection Key Laboratory of Regional Air Quality Monitoring,
22 Guangdong Environmental Monitoring Center, Guangzhou, 510308, China

23 ⁷Guangzhou Hexin Analytical Instrument Limited Company, Guangzhou 510530,
24 China

25
26
27 **Correspondence to:* Mei Li (limei2007@163.com) and Zhen Zhou (zhouzhen@gig.ac.cn)

28 Tel: 86-20-85225991, Fax: 86-20-85225991

29
30
31
32
33
34
35
36
37
38
39
40
41

42 **Abstract:**

43 The formation of oxalic acid and its mixing state in atmospheric particulate
44 matter (PM) were studied using a single particle aerosol mass spectrometer (SPAMS)
45 in the summer and winter of 2014 in Heshan, a supersite in the rural area of the Pearl
46 River Delta (PRD) region in China. Oxalic acid-containing particles accounted for 2.5%
47 and 2.7% in total detected ambient particles in summer and winter, respectively.
48 Oxalic acid was measured in particles classified as elemental carbon (EC), organic
49 carbon (OC), elemental and organic carbon (ECOC), biomass burning (BB), heavy
50 metal (HM), secondary (Sec), sodium-potassium (NaK) and dust. Oxalic acid was
51 found predominantly mixing with sulfate and nitrate during the whole sampling
52 period, likely due to aqueous phase reactions. In summer, oxalic acid-containing
53 particle number and ozone concentration followed a very similar trend, which may
54 reflect the significant contribution of photochemical reactions to oxalic acid formation.
55 The favorable in-situ pH (2-4) and the dominance of transition metal ions in oxalic
56 acid particles can be plausibly explained by the enhanced production of $\bullet\text{OH}$ from
57 Fenton like reaction, which can promote the oxalic acid production from the oxidation
58 of precursors by $\bullet\text{OH}$ in HM type particles. In wintertime, carbonaceous type particles
59 contained a substantial amount of oxalic acid as well as abundant carbon clusters and
60 biomass burning markers. The general existence of nitric acid and high relative acidity
61 ratio in oxalic acid-containing particles indicates an acidic environment during the
62 formation process of oxalic acid. The peak areas of nitrate, sulfate and oxalic had
63 similar temporal change in the carbonaceous type oxalic acid particles, and the
64 organosulfate-containing oxalic acid particles well correlated with total oxalic acid
65 particles during the episode, which suggests the formation of oxalic acid is closely
66 associated with acid-catalyzed reactions of organic precursors.

67

68 **Keywords:** Oxalic acid; Single particles; Mixing state; Photochemical process;
69 Aqueous phase reactions.

70

71 **1. Introduction**

72 Organic aerosol, typically a large fraction of fine particles, contains more than
73 thousands of organic compounds and contributes to visibility reduction,
74 photochemical smog, climate change and adverse health effects (Novakov and Penner,
75 1993;Goldstein and Galbally, 2007;Jimenez et al., 2009;Poschl and Shiraiwa, 2015). A
76 significant component of organic aerosol is secondary organic aerosol (SOA) formed
77 from the gas phase oxidation of volatile organic compounds (VOCs) followed by
78 partitioning of products into particles or from heterogeneous reactions of VOCs with
79 particles (Hallquist et al., 2009;Zhang et al., 2015). Dicarboxylic acids (DCAs) are
80 abundant and ubiquitous constituents in SOA and can be effective tracers for the
81 oxidative processes leading to the formation of SOA (Kawamura and Ikushima,
82 1993;Ervens et al., 2011;Wang et al., 2012;Cheng et al., 2013). DCAs normally have
83 high water solubility and low vapor pressure, so they play important roles in
84 controlling the hygroscopic properties of organic aerosols (Prenni et al., 2003;Ma et
85 al., 2013) and activating cloud condensation nuclei (Booth et al., 2009). The primary
86 emissions of DCAs from anthropogenic sources in urban areas are minor (Huang and
87 Yu, 2007;Stone et al., 2010), and they are mainly derived from secondary oxidation of
88 VOCs and subsequent intermediates (Ho et al., 2010;Myriokefalitakis et al., 2011).
89 High concentrations of DCAs have been observed in biomass burning plume (Kundu
90 et al., 2010;Kawamura et al., 2013) with more than 70% of DCAs produced from
91 photochemical oxidation of water-soluble organic compounds, and only a small
92 contribution from direct biomass burning emission (van Pinxteren et al., 2014).

93 The production of DCAs through photochemical reactions has been reported in
94 many field studies via the analysis of the diurnal and seasonal variations of
95 DCA(Kawamura and Ikushima, 1993;Kawamura and Yasui, 2005;Aggarwal and
96 Kawamura, 2008;Pavuluri et al., 2010;Ho et al., 2011;Wang et al., 2017), but the
97 mechanism of DCAs formation is still not well understood. Oxalic acid is usually the
98 most abundant DCA observed in the field (Kawamura et al., 2004;Ho et al.,
99 2007;Kawamura et al., 2010). A number of ground based and airborne field studies

100 have found a tight correlation between oxalic acid and sulfate in ambient particles and
101 cloud droplets, relating aqueous phase chemistry to the formation of oxalic acid in
102 aerosols and cloud droplets (Yao et al., 2002; Yao et al., 2003; Yu et al.,
103 2005; Sorooshian et al., 2006; Sorooshian et al., 2007a; Sorooshian et al.,
104 2007b; Miyazaki et al., 2009; Wonaschuetz et al., 2012; Wang et al., 2016). In recent
105 years, several model and laboratory studies suggested that the aqueous phase
106 oxidation of highly water-soluble organics like glyoxal, methylglyoxal and glyoxylic
107 acid can efficiently produce oxalic acid in aerosol particles and cloud droplets (Lim et
108 al., 2010; Myriokefalitakis et al., 2011; Ervens et al., 2014; Yu et al., 2014; McNeill,
109 2015). Recent stable carbon isotope studies and field observations have also suggested
110 that oxalic acid forms through aqueous phase reactions (Wang et al., 2012; Cheng et al.,
111 2015). However, the detailed formation mechanisms of oxalic acid from
112 photochemistry and aqueous phase chemistry in ambient aerosols are still not
113 comprehensively understood and need to be further studied.

114 Online measurements of the size distribution of oxalic acid-containing particles
115 and the mixing state of oxalic acid with other compounds in aerosols are useful to
116 examine the formation and evolution of oxalic acid and SOA particles. Sullivan and
117 Prather investigated the diurnal cycle and mixing state of DCA-containing particles in
118 Asian aerosol outflow using aerosol time-of-flight mass spectrometry (ATOFMS), and
119 proposed the formation of DCA on Asian dust (Sullivan and Prather, 2007b). In
120 addition, Yang et al. (2009) measured oxalic acid particles in Shanghai and proposed
121 that in-cloud processes and heterogeneous reactions on hydrated aerosols contributed
122 to the formation of oxalic acid (Yang et al., 2009). While the formation mechanism of
123 oxalic acid especially in urban areas is still not clear, online measurements of the
124 mixing state of oxalic acid provide a powerful tool to better understand the formation
125 of oxalic acid in aerosol particles and cloud droplets.

126 The Pearl River Delta (PRD) region has distinct meteorological seasonality
127 under the influence of the Asian monsoon system, which brings air from the ocean in
128 spring and summer, and carries polluted air from northern China in autumn and winter.
129 Strong photochemical activity occurs in summer under the condition of high

130 temperature and relative humidity, and in winter high loadings of particles from
131 northern cities are favorable for the occurrence of haze episode (Bi et al., 2011;Zhang
132 et al., 2013;Zhang et al., 2014). Here we present the seasonal field measurements of
133 the mixing state of oxalic acid-containing particles using a single particle aerosol
134 mass spectrometer (SPAMS) in a rural supersite of the PRD region. The seasonal
135 characteristic of oxalic acid particles and mixing state with secondary species were
136 investigated to explore the formation mechanisms of oxalic acid and aging process of
137 SOA.

138 **2. Methods**

139 **2.1 Aerosol sampling**

140 Particles were sampled using a single particle aerosol mass spectrometer
141 (SPAMS) at the Guangdong Atmospheric Supersite (22.73N, 112.93E), a rural site at
142 Heshan city (Figure S1). The supersite is surrounded by farm land and villages, with
143 no local industrial or traffic emissions. Ambient aerosols were sampled to the SPAMS
144 through a 2.5m long copper tube with 0.5m of the sampling inlet located above the top
145 of the building. The measurement period was from July 18 to August 1 in 2014, and
146 from January 27 to February 8 in 2015. Real-time PM_{2.5} mass concentration was
147 simultaneously measured by a TEOM monitor (series 1405, Thermo scientific), and
148 hourly concentrations of O₃ were measured by an O₃ analyzer (model 49i, Thermo
149 scientific). The local meteorological data including temperature, relative humidity and
150 visibility were measured on the rooftop of the building. The average temperature
151 during the field study was 29.5 °C in summer and 14.1 °C in winter and the average
152 relative humidity was 71.7% and 63% in summer and winter, respectively.

153 **2.2 SPAMS**

154 Real-time measurements of single atmospheric particles has been demonstrated
155 by Prather and co-workers in the 1990s using aerosol time-of-flight mass
156 spectrometry (ATOFMS) (Prather et al., 1994;Noble and Prather, 1996). Based on the
157 same principle, the single particle aerosol mass spectrometer (SPAMS) developed by
158 Guangzhou Hexin Analytical Company was applied to field measurements of single

159 particles in the current work. The details of the SPAMS system have been introduced
160 previously (Li et al., 2011). Briefly, aerosol particles are sampled into the vacuum
161 pumped aerodynamic lens of the SPAMS through an electro-spark machined 80 μ m
162 critical orifice at a flow rate of 75 ml min⁻¹. The individual particles with a terminal
163 velocity are introduced to the sizing region. The velocity of each single particle is
164 detected by two continuous laser beams (diode Nd:YAG, 532 nm) with a space of 6
165 cm. The velocity is then used to calculate the single particle aerodynamic diameter
166 and provide the precise timing of the firing of a 266 nm laser used to induce
167 desorption and ionization (Nd:YAG laser, 266nm). The energy of the
168 desorption/ionization 266 nm laser was 0.6 mJ and the power density was kept at
169 about 1.6 \times 10⁸ W/cm² during both sampling periods. The 266 nm laser generates
170 positive and negative ions that are detected by a Z-shaped bipolar time of flight mass
171 spectrometer. The size range of the detected single particles is 0.2 to 2 μ m.
172 Polystyrene latex spheres (Nanosphere size standards, Duke Scientific Corp., Palo
173 Alto) of 0.22-2.0 μ m diameter were used for size calibration.

174 **2.3 Data analysis**

175 The size and chemical composition of single particles detected by SPAMS were
176 analyzed using the COCO toolkit based on the Matlab software. Particles were
177 clustered into several groups using the neural network algorithm (ART-2a) to group
178 particles into clusters with similar mass spectrum features. The ART-2a parameters
179 used in this work were set to a vigilance factor of 0.8, a learning rate of 0.05, and a
180 maximum of 20 iterations. We collected 516,679 and 767,986 particles with both
181 positive and negative mass spectra in summer and winter respectively. A standard
182 solution of oxalic acid was prepared with pure oxalic acid (H₂C₂O₄, purity: 99.99%,
183 Aladdin Industrial Corporation) and atomized to aerosols. After drying through two
184 silica gel diffusion driers, pure oxalic acid particles were directly introduced into the
185 SPAMS. The positive and negative mass spectra of oxalic acid are shown in Figure S2.
186 Based on the mass spectra of pure oxalic acid and previous ambient measurements by
187 ATOFMS (Silva and Prather, 2000; Sullivan and Prather, 2007b; Yang et al., 2009),
188 HC₂O₄⁻ (m/z -89) is selected as the ion peak for oxalic acid containing particles. In

189 this work, oxalic acid particles are identified if the peak area of m/z -89 was larger
190 than 0.5% of the total signal in the mass spectrum. With this threshold, 13109 and
191 20504 of oxalic acid-containing particles were obtained in summer and winter
192 separately, accounting for 2.5% and 2.7% of the total detected particles. The
193 percentage of oxalic acid-containing particles in total particles in this work was
194 comparable to the reported value in the urban area of Shanghai (3.4%) (Yang et al.,
195 2009). However, these percentages are in general much lower than those reported in
196 cleaner environments such as the western Pacific Ocean where oxalic acid was found
197 in up to 1-40% of total particles due to little anthropogenic influences (Sullivan and
198 Prather, 2007b).

199 The oxalic acid containing particles are classified into eight types in the
200 following order: elemental carbon (EC), organic carbon (OC), elemental and organic
201 carbon (ECOC), biomass burning (BB), heavy metal (HM), secondary (Sec),
202 sodium-potassium (NaK) and dust. Different type particles are identified according to
203 characteristic ion markers and dominant chemical species (Table S1): (1) particles
204 containing abundant carbon clusters like $\pm 12[\text{C}]^{+/-}$, $\pm 24[\text{C}_2]^{+/-}$, $\pm 36[\text{C}_3]^{+/-}$ with relative
205 peak area more than 0.5% are classified as EC type, (2) any remaining particles
206 containing abundant signals of $27[\text{C}_2\text{H}_3]^+$, $43[\text{C}_2\text{H}_3\text{O}]^+$ and hydrocarbon clusters with
207 relative peak area more than 0.5% are classified as OC type, (3) any remaining
208 particles containing signals of $\pm 12[\text{C}]^{+/-}$, $\pm 24[\text{C}_2]^{+/-}$, $37[\text{C}_3\text{H}]^+$ and $43[\text{C}_2\text{H}_3\text{O}]^+$ with
209 relative peak area more than 0.5% are classified as ECOC type, (4) any remaining
210 particles containing abundant signals of $39[\text{K}]^+$ (peak area > 1500) with relative peak
211 area of $-59[\text{C}_2\text{H}_3\text{O}_2]^-$ and $-73[\text{C}_3\text{H}_5\text{O}_2]^-$ simultaneously more than 0.5% are classified
212 as BB type, (5) any remaining particles containing signals of $55[\text{Mn}]^+$, $56[\text{Fe}]^+$,
213 $63/65[\text{Cu}]^+$, $64[\text{Zn}]^+$ and $208[\text{Pb}]^+$ with relative peak area more than 0.5% are
214 classified as HM type, (6) any remaining particles containing abundant signals of
215 $18[\text{NH}_4]^+$ (peak area > 50), $-62[\text{NO}_3]^-$ (peak area > 100) and $-97[\text{HSO}_4]^-$ (peak area > 100)
216 are classified as Sec type, (7) any remaining particles containing abundant signals of
217 $23[\text{Na}]^+$ (peak area > 1500) and related species are classified as NaK type, (8) any
218 remaining particles containing signals of $40[\text{Ca}]^+$, $56[\text{CaO}]^+$ and related species are

219 classified as dust type. The rules for oxalic acid particles classification in the current
220 work have been reported in previous studies (Sullivan and Prather, 2007a; Yang et al.,
221 2009; Zhang et al., 2013; Li et al., 2014).

222 2.4 Inorganic ions and in-situ pH (pH_{is})

223 Water-soluble inorganic ions and trace gases were determined by an online
224 analyzer for monitoring aerosols and gases (MARGA, model ADI 2080, Applikon
225 Analytical B. V. Corp., the Netherlands) with a $PM_{2.5}$ sampling inlet at one hour
226 resolution from July 18 to August 1 in 2014. The principle and instrumental design
227 has been described in detail elsewhere (ten Brink et al., 2007; Du et al., 2011; Behera et
228 al., 2013; Khezri et al., 2013). Standard solutions containing all detected ions were
229 injected into MARGA before and after the field measurement. The liquid water
230 content and the concentration of H^+ in particles are calculated using the ISORROPIA
231 II model (Nenes et al., 1998, 1999; Fountoukis and Nenes, 2007). We choose stable
232 mode and reverse type in the ISORROPIA model to calculate the concentration of H^+
233 and the liquid water content in this work. The in-situ pH (pH_{is}) of particles is
234 calculated through the following equation:

$$235 \quad pH_{is} = -\log \alpha_{H^+} = -\log(\gamma_{H^+} \times n_{H^+} \times 1000/V_a) \quad (1)$$

236 where n_{H^+} is the concentration of H^+ (mol m^{-3}) and V_a is the volume concentration of
237 the H_2O ($\text{cm}^3 \text{m}^{-3}$), while γ_{H^+} is the activity coefficient of H^+ (Xue et al., 2011; Cheng
238 et al., 2015). The temporal variation of pH_{is} of ambient $PM_{2.5}$ particles is presented in
239 Figure S3, and demonstrated that 97% of particles were acidic in summer.

240 3. Results and Discussion

241 3.1 Seasonal variation of oxalic acid containing particles

242 The clustered 48 hr back trajectories of air masses arriving in Heshan during the
243 sampling period are shown in Figure S4. In summer, air masses at 500m levels above
244 the ground were mainly from the ocean and rural areas with less influence of human
245 activity, while in winter air masses were directly from urban areas of Guangzhou and
246 Foshan, indicating a strong influence from anthropogenic emissions. The temporal

247 variations of the total detected particles and oxalic acid containing particles in
248 summer and winter are shown in Figure 1. The total particles had similar trends with
249 the mass concentration of ambient PM_{2.5}, suggesting that the counts of total particles
250 detected by SPAMS can be representative of PM_{2.5} mass concentration during the
251 whole sampling periods. The oxalic acid (C₂-containing) particles, in general,
252 exhibited distinct diurnal peaks from July 28 to August 1, while they showed different
253 temporal trends in winter. The relative abundance of oxalic acid particles in all of the
254 sampled particles (C₂/total ratio) had the same variation with the abundance of oxalic
255 acid particles in summer, especially in the period of July 28 – August 1 (Figure 1). In
256 winter, however, particle counts and relative abundance of oxalic acid had different
257 temporal changes except Jan 30 and February 5-8, when the count and relative
258 abundance of oxalic acid particles simultaneously had a sudden increase.

259 The oxalic acid-containing particles were clustered into eight groups, and they
260 altogether accounted for 89.6% and 95.1% of total oxalic acid particles in summer and
261 winter, respectively. Table 1 shows that in summer heavy metal (HM) type particles
262 contributed 31.3% to total oxalic acid particles, followed by the Sec (19.2%) and BB
263 type (13%). However, in winter BB type particles were the most abundant and
264 accounted for 24.2% of the oxalic acid-containing particles, followed by EC and HM
265 type. Besides, carbonaceous type particles including EC, OC, ECOC and BB
266 accounted for 28.1% of oxalic acid particles in summer and 59.8% in winter,
267 indicating the seasonal different characteristics of oxalic acid particles. The temporal
268 variations of eight groups of oxalic acid particles in summer and winter are illustrated
269 in Figure 1. In summer HM type particles (orange color) and total oxalic acid particles
270 exhibited similar diurnal patterns, suggesting a possible connection between the
271 production of oxalic acid and the transition metals (e.g. Fe, Cu) (Zhou et al., 2015).
272 Although Sec, BB and EC type particles showed similar diurnal patterns with total
273 oxalic acid particles, the concentrations of these type particles were generally lower
274 than HM type particles. In winter diurnal variation of oxalic acid particles was not
275 obvious but a sharp increase, accompanied by the increase of BB, EC and Sec type
276 particles, was observed on February 8.

277 The averaged positive and negative ion mass spectra of oxalic acid containing
278 particles are shown in Figure 2. The positive ion spectrum of oxalic acid particles in
279 summer was characterized by high fractions of metal ion peaks including $23[\text{Na}]^+$,
280 $27[\text{Al}]^+$, $39[\text{K}]^+$, $55[\text{Mn}]^+$, $56[\text{Fe}]^+$, $63/65[\text{Cu}]^+$, $64[\text{Zn}]^+$ and $208[\text{Pb}]^+$, and
281 carbonaceous marker ions at m/z $27[\text{C}_2\text{H}_3]^+$, $36[\text{C}_3]^+$, $43[\text{C}_2\text{H}_3\text{O}/\text{C}_3\text{H}_7]^+$, $48[\text{C}_4]^+$
282 (Figure 2 a). The negative ion spectrum of oxalic acid particles in summer was
283 characterized by the strong intensity of secondary ions including m/z $-46[\text{NO}_2]^-$,
284 $-62[\text{NO}_3]^-$, $-79[\text{PO}_3]^-$, $-80[\text{SO}_3]^-$, $-96[\text{SO}_4]^-$ and $-97[\text{HSO}_4]^-$, as well as carbon clusters
285 of $-24[\text{C}_2]^-$, $-36[\text{C}_3]^-$, $-48[\text{C}_4]^-$ and BB markers of $-59[\text{C}_2\text{H}_3\text{O}_2]^-$ and $-73[\text{C}_3\text{H}_5\text{O}_2]^-$
286 (Figure 2 b) (Zauscher et al., 2013). More carbonaceous clusters, i.e., $27[\text{C}_2\text{H}_3]^+$,
287 $29[\text{C}_2\text{H}_5]^+$, $36[\text{C}_3]^+$, $37[\text{C}_3\text{H}]^+$, $43[\text{C}_2\text{H}_3\text{O}]^+$, $48[\text{C}_4]^+$, $51[\text{C}_4\text{H}_3]^+$, $55[\text{C}_4\text{H}_7]^+$, $60[\text{C}_5]^+$,
288 $63[\text{C}_5\text{H}_3]^+$, $65[\text{C}_5\text{H}_5]^+$, $74[\text{C}_2\text{H}_2\text{O}_3]^+$, $77[\text{C}_6\text{H}_5]^+$, were observed in the positive ion
289 spectrum of oxalic acid particles in winter (Figure 2 c) than in summer. The negative
290 ion spectrum of oxalic acid particles in winter (Figure 2 d) contained a large amount
291 of secondary ions, similar to those found in summer, and a more intense signal of
292 nitric acid ($-125[\text{HNO}_3\text{NO}_3]^-$), suggesting an acidic nature of oxalic acid particles in
293 winter.

294 The mixing state of oxalic acid particles with sulfate, nitrate and ammonium
295 (SNA) was investigated through the percentage of SNA-containing oxalic acid
296 particles in total oxalic acid particles (Figure 3). Oxalic acid was found to be
297 internally mixed with sulfate and nitrate during both sampling periods with percentage
298 of 93% and 94% in summer respectively, and both 98% in winter (Figure 3 a).
299 However, the percentage of NH_4^+ with oxalic acid was only 18% in summer but
300 increased to 71% in winter. Linear correlations between NH_4^+ -containing oxalic acid
301 particles ($\text{C}_2\text{-NH}_4^+$) and total oxalic acid particles are depicted in Figure 3, with better
302 linear regression ($r^2=0.98$) in winter than summer. The low percentage of NH_4^+ in
303 oxalic acid particles in summer indicated the presence of oxalic acid in NH_4^+ -poor
304 particles. Aqueous phase production of SO_4^{2-} has been studied well and the linear
305 correlation between oxalic acid and SO_4^{2-} has been used to study the production of
306 oxalic acid through aqueous phase reactions (Yu et al., 2005; Miyazaki et al.,

307 2009;Cheng et al., 2015). In our work, oxalic acid and $C_2-SO_4^{2-}$ displayed good
308 correlations in summer and winter (both $r^2=0.99$), which suggests a common
309 production route of oxalic acid and sulfate, likely aqueous phase reactions.

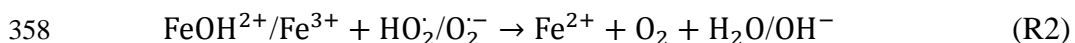
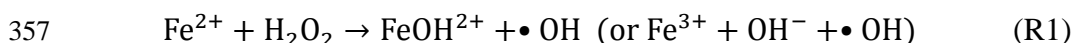
310 Figure 4 shows the unscaled size-resolved number distributions of the eight types
311 of oxalic acid particles. Oxalic acid mainly existed in 0.4 to 1.2 μm particles during
312 the entire sampling period but exhibited different peak modes for each particle type in
313 summer and winter. In summer, major types of oxalic acid particles showed distinct
314 peak mode at different size diameter. EC and Sec type particles peaked at 0.5 μm ,
315 followed by BB type particles at 0.55 μm , then HM type particles at 0.6 μm , and OC
316 type particles at 0.7 μm . The difference of peak mode suggests the possible different
317 chemical evolution process for each type oxalic acid-containing particles. However, in
318 winter, oxalic acid particles showed broader size distribution from 0.5 to 0.8 μm for
319 all particle types. Oxalic acid particles of all types were generally larger in winter than
320 summer, possibly due to condensation and coagulation of particles during aging of
321 oxalic acid particles in winter.

322 **3.2 Photochemical production of oxalic acid in summer**

323 In summer oxalic acid particles showed peaks in the afternoon especially from
324 July 28 to August 1, which was in agreement with the variation pattern of the O_3
325 concentration (Figure 5), indicating a strong association of oxalic acid formation with
326 photochemical reactions. Malonic acid is another product of photochemical oxidation
327 of organic compounds (Kawamura and Ikushima, 1993;Wang et al., 2012;Meng et al.,
328 2013;Meng et al., 2014). In our campaign, malonic acid containing particles had
329 diurnal trends similar to oxalic acid particles and O_3 concentration. As the dominant
330 particle type, HM particles had identical variation pattern with total oxalic acid
331 particles. They are characterized by highly abundant metal ion peaks like 55[Mn]⁺,
332 56[Fe]⁺, 63/65[Cu]⁺, 64[Zn]⁺ and 208[Pb]⁺, as well as secondary ion peaks of
333 -46[NO₂]⁻, -62[NO₃]⁻, -80[SO₃]⁻, -96[SO₄]⁻ and -97[HSO₄]⁻ in the negative spectrum
334 in summer (Figure 6). •OH produced from Fenton reactions between H₂O₂ and Fe^{2+/3+}
335 in acidic solutions has been considered as a substantial source of •OH(Fenton,
336 1894;Dunford, 2002;Herrmann et al., 2015). The high abundance of metal ions in

337 oxalic acid particles may be an indication of possible Fenton reactions in the acidic
338 aqueous phase of acidic particles (pH<5, Figure S3), although we cannot exclude the
339 possibility of gas phase condensation of oxalic onto HM particles.

340 The oxidation of glyoxal and glyoxylic acid by •OH has been identified as an
341 important pathway of oxalic acid production by field and laboratory studies (Ervens et
342 al., 2004;Ervens and Volkamer, 2010;Wang et al., 2012;Wang et al., 2015). The
343 modeling studies from Ervens et al. (2014) suggest that oxalic acid production from
344 glyoxal and glyoxylic acid in aqueous phase significantly depends on •OH availability
345 (Ervens et al., 2014). The main sources of •OH in the aqueous phase contain both
346 direct uptake from the gas phase and the chemical sources in the aqueous phase such
347 as Fenton type reactions and photolysis of H₂O₂, NO₃⁻, NO₂⁻, and chromophoric
348 dissolved organic matter (CDOM) (Yu et al., 2014;Badali et al., 2015;Ervens,
349 2015;Gligorovski et al., 2015;Herrmann et al., 2015;Tong et al., 2016). Considering
350 the low Henry's law constant of •OH (K_{H,OH}=30 M atm⁻¹) (Hanson et al., 1992) and
351 abundant fraction of transition metal ions in the oxalic acid particles, the photolysis of
352 H₂O₂ through Fenton reactions involving the catalysis of transition metal ions like
353 Fe^{2+/3+}, Cu⁺²⁺ and Mn^{2+/3+} likely contributes substantially to the source of •OH in the
354 aqueous phase in this work (Deguillaume et al., 2005;Herrmann et al., 2005;Ervens et
355 al., 2014). The •OH formation process through Fenton reactions can be expressed as
356 (Ervens, 2015):



359 The actual chemical process is far more complex and involves iron oxides and
360 iron-complexes, thus in the current work we focus on the potential availability of •OH
361 from Fenton reactions and the impact on the oxidation process of organic precursors.

362 In order to investigate the photochemical aqueous phase formation of oxalic acid
363 in summer, the diurnal variations of O₃, oxalic acid particles, HM group particles and
364 *pH_{is}* of ambient particles averaged from July 28 to August 1, 2014 are shown in
365 Figure 7. The concentration of O₃ increased after 9:00 and peaked at 17:00, while
366 oxalic acid particles and HM group particles both increased after 10:00 and showed

367 two peaks at 15:00 and 19:00. The prominent photochemical feature of oxalic acid
368 particles suggested a close association of photochemical reactions with oxalic acid
369 production. Although $\bullet\text{OH}$ production from Fenton reactions can both occur under
370 dark and light radiation conditions, photo-Fenton reactions may have more
371 contribution to the enhancement of oxalic acid particles in the current work. This was
372 possibly due to the diurnal variation of pH_{is} , since Fenton reactions strongly depend
373 on the pH of the aqueous phase (Gligorovski et al., 2015). When $\text{pH} < 1$, Fe^{2+} is
374 directly oxidized by H_2O_2 to Fe^{3+} with no production of $\bullet\text{OH}$ (Barb et al.,
375 1951; Kremer, 2003), and the most favorable pH value for Fenton reaction is between
376 2.5 and 5 (Deguillaume et al., 2005). In the current work the pH_{is} of ambient particles
377 ranged from -1.42 to 4.01, and the influences of pH_{is} from RH and inorganic ions are
378 discussed in Figure S5. Strongly acidic particles were observed during the whole day
379 with high pH_{is} at 6:00 and after 12:00. Although pH_{is} was around 2 at 6:00, only a few
380 oxalic acid-containing particles were observed during this period due to low
381 abundance of HM particles. Oxalic acid-containing particles were found to increase
382 from 12:00 to 21:00, which was attributed to increased organic precursors from VOCs
383 oxidation and enhanced $\bullet\text{OH}$ production from Fenton reactions under pH_{is} at 1-4. The
384 number concentration of oxalic acid particles peaked at 19:00 instead of during the
385 strong photochemical activity period in the afternoon; this was possibly due to the
386 efficient degradation of oxalic acid from the complex with iron (Sorooshian et al.,
387 2013; Zhou et al., 2015). Furthermore, photolysis of $\text{Fe}(\text{oxalate})_n^{3-2n}$ can contribute to
388 99% of the overall degradation of oxalic acid (Weller et al., 2014). Although the
389 enhanced $\bullet\text{OH}$ production from photo-Fenton reactions was favorable for the
390 formation of oxalic acid from 12:00 to 18:00, we speculate that a high degradation
391 rate of oxalic acid by iron complexation resulted in a lower net production of oxalic
392 acid than at 19:00. In addition to the contribution from Fenton reactions after 12:00,
393 the precursors of oxalic acid such as glyoxylic acid have higher reaction rate with
394 $\bullet\text{OH}$ in high pH solutions due to reported studies (Ervens et al., 2003; Herrmann,
395 2003; Cheng et al., 2015), thus the increase of pH not only enhances $\bullet\text{OH}$ production
396 from photo-Fenton reactions, but also promotes the oxidation process of oxalic acid's

397 precursors by •OH.

398 Based on above discussions of the mixing state of oxalic acid with secondary
399 ions and transition metals, a plausible explanation to the formation process of oxalic
400 acid in the HM type oxalic acid particles is proposed (Figure S6). In summer strong
401 photochemical activity and high O₃ concentrations in the afternoon lead to more
402 production of reactive radicals such as •OH and HO₂•, which promote the oxidation of
403 VOCs to dicarbonyls and aldehydes (e.g. glyoxal and methylglyoxal), followed by a
404 subsequent partitioning into the aqueous phase of particles (Myriokefalitakis et al.,
405 2011). Acidic particles containing transition metals like Fe and Cu potentially yield
406 more •OH in acidic aqueous phase, then hydrated dicarbonyls and aldehydes can be
407 oxidized by •OH to glyoxylic acid and finally to oxalic acid (Wang et al., 2012)
408 (Figure S6). Recently Ma et al. (2015) had studied the Fe-containing particles in the
409 PRD and found Fe-containing particles are more efficient at generating •OH in
410 summer than winter (Ma et al., 2015), which supports the enhanced •OH production
411 in HM type particles in this work. A large amount of Fe related particles are emitted
412 from steel industries in the North China Plain and metals like V, Zn, Cu and Pb from
413 electronic manufacturing (Cui and Zhang, 2008; Dall'Osto et al., 2008). These metals
414 contribute significantly to haze episodes (Moffet et al., 2008; Li et al., 2014), which
415 possibly increases the formation of SOA by yielding more •OH participating the
416 heterogeneous and aqueous reactions.

417 The diurnal patterns of temperature, wind speed are presented in Figure S7. The
418 high temperature between 9:00 and 19:00 was favorable to the secondary processing
419 of organic precursors. The wind speed was low during the whole day, especially
420 between 9:00 and 18:00, which provided a stagnant environment for the increase in
421 oxalic acid produced from photochemical process. The influence from traffic emission
422 was investigated through the diurnal variations of total EC type particles and NO₂
423 (Figure S7). The EC type particles increased from 12:00 to 21:00, which had same
424 variation as total oxalic acid, but NO₂ followed the rush hour pattern with two peaks
425 from 5:00 to 8:00 and from 18:00 to 21:00. Traffic emission is not expected to have a
426 large contribution to oxalic acid in this study.

427 3.3 Formation process of oxalic acid in winter

428 Despite lower O₃ concentrations and photochemical activity in winter, oxalic
429 acid particles were still prevalent in carbonaceous particles, especially BB type
430 particles. While oxalic acid was found to be internally mixed with sulfate and nitrate
431 both in summer and winter, the nitric acid was only observed in oxalic acid particles
432 in winter, indicating a strongly acidic nature of oxalic acid particles in winter.
433 Considering a possible connection of oxalic acid production with the acidic
434 environment, the temporal concentrations of oxalic acid, sulfate and nitrate were
435 investigated through their peak areas in the carbonaceous type oxalic acid particles
436 including EC, OC, ECOC and BB type in Figure 8. The peaks of m/z -62[NO₃]⁻ and -
437 97[HSO₄]⁻ represent nitrate and sulfate, respectively. Nitrate, sulfate and oxalic acid
438 showed very similar variation patterns in winter, indicating a close connection of their
439 co-existence. Although nitric acid was found in the oxalic acid particles, the acidity of
440 the oxalic acid particles was not estimated since the real-time concentration of
441 inorganic ions was not available during the sampling period in winter. Instead the
442 relative acidity ratio (R_{ra}), defined as the ratio of total peak areas of nitrate and sulfate
443 to the peak area of ammonium (m/z 18[NH₄]⁺), was used (Denkenberger et al.,
444 2007;Pratt et al., 2009). The R_{ra} of carbonaceous type oxalic acid particles ranged
445 from 7 to 114 with an average value of 25 (Figure 8), indicating an intensely acidic
446 environment of carbonaceous type oxalic acid particles in winter. Several studies have
447 reported potential production of oxalic acid from acid-catalyzed aqueous phase
448 reactions in aerosols (Carlton et al., 2006;Carlton et al., 2007;Tan et al., 2009). In this
449 work the acidic environment of the carbonaceous type oxalic acid particles and
450 similar variation patterns among oxalic acid, sulfate and nitrate may suggest a
451 relationship between the degradation of organic precursors and the acidic chemical
452 process. However, the temporal change of R_{ra} did not follow a similar trend as the
453 peak area of oxalic acid in most particles, possibly due to the multi-step formation of
454 oxalic acid influenced by many factors such as precursors, liquid water content and
455 ion strength (Carlton et al., 2007;Cheng et al., 2013;Cheng et al., 2015).

456 The sharp increase of oxalic acid particles on February 8, 2015 (Figure 1) was

457 selected as a typical episode to investigate the formation processes of oxalic acid in
458 winter. During the episode, the 48 hr back trajectory analysis showed air masses that
459 originated from the urban areas of Guangzhou and Foshan city (Figure S4), indicating
460 strong influence on organic precursors from anthropogenic emissions. Oxalic acid
461 particle types were dominated by BB (23.2%), followed by EC (22.0%) and Sec
462 (15.1%) type (Table 2). Carbonaceous particles including EC, ECOC, OC, BB
463 accounted for 61.6% of the total oxalic acid particles. The mass spectra of oxalic acid
464 particles were characterized by many hydrocarbon clusters of $27[\text{C}_2\text{H}_3]^+$, $29[\text{C}_2\text{H}_5]^+$,
465 $37[\text{C}_3\text{H}]^+$, $43[\text{C}_2\text{H}_3\text{O}]^+$, $51[\text{C}_4\text{H}_3]^+$, $55[\text{C}_4\text{H}_7]^+$, $63[\text{C}_5\text{H}_3]^+$, $65[\text{C}_5\text{H}_5]^+$, $74[\text{C}_2\text{H}_2\text{O}_3]^+$,
466 $77[\text{C}_6\text{H}_5]^+$, and carbon clusters of $36[\text{C}_3]^+$, $48[\text{C}_4]^+$, $60[\text{C}_5]^+$ in positive mass spectrum,
467 while the negative mass spectrum was characterized by elemental carbon clusters like
468 $-24[\text{C}_2]^-$, $-36[\text{C}_3]^-$, $-48[\text{C}_4]^-$, biomass burning markers of $-59[\text{C}_2\text{H}_3\text{O}_2]^-$ and
469 $-73[\text{C}_3\text{H}_5\text{O}_2]^-$ and secondary species including $-42[\text{CNO}]^-$, $-46[\text{NO}_2]^-$, $-62[\text{NO}_3]^-$,
470 $-79[\text{PO}_3]^-$, $-80[\text{SO}_3]^-$, $-96[\text{SO}_4]^-$ and $-97[\text{HSO}_4]^-$ (Figure 9 a).

471 As the precursor of oxalic acid, glyoxal has the potential to react with sulfuric
472 acid to produce organosulfates through acid-catalyzed nucleophilic addition according
473 to laboratory and chamber studies (Surratt et al., 2007; Galloway et al., 2009). The
474 negative ion of $-155([\text{C}_2\text{H}_3\text{O}_2\text{SO}_4]^-)$ has been identified as the marker ion of
475 organosulfates derived from glyoxal in chamber and field measurements using
476 ATOFMS (Surratt et al., 2008; Hatch et al., 2011). The formation of organosulfates
477 from glyoxal requires an acidic aqueous environment, which can be used as an
478 indicator of acid-catalyzed ageing process of organic compounds. The temporal trend
479 of organosulfate-containing oxalic acid particles in winter is shown in Figure S8,
480 which exhibited a similar pattern as the total oxalic acid particles during the whole
481 sampling period in winter. The percentage of organosulfate-containing oxalic acid
482 particles in total oxalic acid particles ranged from 0 to 16.4% with the highest ratio
483 observed in the episode (February 8). The linear regression between total oxalic acid
484 particles and organosulfate-containing oxalic acid particles in the episode is exhibited
485 in Figure 9b, and the robust correlation ($r^2=0.81$) between them suggests that oxalic
486 acid and organosulfate may share similar formation process. Based on the above

487 discussion, the degradation of carbonaceous species associated with acid-catalyzed
488 reactions may have an important contribution to the formation of oxalic acid during
489 the episode in winter. Similar particle types and mass spectra of oxalic acid-containing
490 particles during the episode and the whole sampling period in winter were observed,
491 which suggest the acid-catalyzed oxidation of organic precursors as a potential source
492 for oxalic acid.

493 **4. Summary and conclusions**

494 Oxalic acid containing particles were measured by a single particle aerosol mass
495 spectrometer (SPAMS) in the summer and winter of 2014 in Heshan, China. They
496 accounted for 2.5% and 2.7% of the total detected ambient particles. In summer heavy
497 metal-containing particles were the largest group of particles containing oxalic acid
498 with a fraction of 31.3% followed by Sec type (19.2%), while in winter BB type was
499 the dominant group with a percentage of 24.2%. More than 90% of oxalic acid
500 particles were internally mixed with sulfate and nitrate during the whole sampling
501 period. Only 18% of oxalic acid particles contained ammonium in summer, which
502 increased to 71% in winter. In summer oxalic acid and O₃ concentration exhibited
503 similar diurnal variations, indicating a substantial contribution of photochemical
504 reactions to oxalic acid formation. Furthermore, suitable in-situ pH is favorable for
505 Fenton like reactions to produce •OH in HM type particles, and might promote the
506 oxalic acid production from the oxidation of precursors by •OH in HM type particles.
507 In winter carbonaceous type particles including EC, OC, ECOC and BB groups
508 accounted for 59.8% of oxalic acid particles and increased to 61.6% in the episode.
509 Nitric acid and organosulfate were found to co-exist in oxalic acid-containing
510 particles in the winter, which suggests a close association with acid-catalyzed
511 reactions. Acid-catalyzed oxidation of organic precursors is a potential contribution
512 for the formation of oxalic acid in winter. The current study demonstrates that SPAMS
513 is a unique tool for understanding the mixing states of different components of
514 ambient aerosols, which are useful for exploring the formation and evolution process
515 of SOA.

516 **Acknowledgements**

517 This work was financially supported by National Key Technology R&D Program
518 (Grant No. 2014BAC21B01), Guangdong Province Public Interest Research and
519 Capacity Building Special Fund (Grant No. 2014B020216005), the Strategic Priority
520 Research Program (B) of the Chinese Academy of Sciences (Grant No.
521 XDB05040502), Guangdong Industry-University Research Program (Grant
522 No.2012B090500014), and NSFC of Guangdong Province (Grant No.
523 2015A030313339). Chak K. Chan would like to acknowledge funding support of the
524 General Fund of National Natural Science Foundation of China (Grant No. 41675117).
525 Haijie Tong acknowledge Max Planck Society for funding and Ulrich Pöschl for
526 helpful discussions.

527 **References**

- 528 Aggarwal, S. G., and Kawamura, K.: Molecular distributions and stable carbon
529 isotopic compositions of dicarboxylic acids and related compounds in aerosols
530 from Sapporo, Japan: Implications for photochemical aging during long-range
531 atmospheric transport, *Journal of Geophysical Research-Atmospheres*, 113,
532 D14301, 10.1029/2007jd009365, 2008.
- 533 Badali, K. M., Zhou, S., Aljawhary, D., Antinolo, M., Chen, W. J., Lok, A., Mungall,
534 E., Wong, J. P. S., Zhao, R., and Abbatt, J. P. D.: Formation of hydroxyl radicals
535 from photolysis of secondary organic aerosol material, *Atmospheric Chemistry
536 and Physics*, 15, 7831-7840, 10.5194/acp-15-7831-2015, 2015.
- 537 Barb, W., Baxendale, J., George, P., and Hargrave, K.: Reactions of ferrous and ferric
538 ions with hydrogen peroxide. Part I.—The ferrous ion reaction, *Transactions of
539 the Faraday Society*, 47, 462-500, 1951.
- 540 Behera, S. N., Betha, R., Liu, P., and Balasubramanian, R.: A study of diurnal
541 variations of PM 2.5 acidity and related chemical species using a new
542 thermodynamic equilibrium model, *Science of the Total Environment*, 452,
543 286-295, 2013.
- 544 Bi, X., Zhang, G., Li, L., Wang, X., Li, M., Sheng, G., Fu, J., and Zhou, Z.: Mixing
545 state of biomass burning particles by single particle aerosol mass spectrometer in
546 the urban area of PRD, China, *Atmospheric Environment*, 45, 3447-3453, 2011.
- 547 Booth, A. M., Topping, D. O., McFiggans, G., and Percival, C. J.: Surface tension of
548 mixed inorganic and dicarboxylic acid aqueous solutions at 298.15 K and their
549 importance for cloud activation predictions, *Phys Chem Chem Phys*, 11,
550 8021-8028, 10.1039/b906849j, 2009.
- 551 Carlton, A. G., Turpin, B. J., Lim, H. J., Altieri, K. E., and Seitzinger, S.: Link
552 between isoprene and secondary organic aerosol (SOA): Pyruvic acid oxidation

553 yields low volatility organic acids in clouds, *Geophysical Research Letters*, 33,
554 L06822, 10.1029/2005gl025374, 2006.

555 Carlton, A. G., Turpin, B. J., Altieri, K. E., Seitzinger, S., Reff, A., Lim, H. J., and
556 Ervens, B.: Atmospheric oxalic acid and SOA production from glyoxal: Results
557 of aqueous photooxidation experiments, *Atmospheric Environment*, 41,
558 7588-7602, 10.1016/j.atmosenv.2007.05.035, 2007.

559 Cheng, C., Wang, G., Meng, J., Wang, Q., Cao, J., Li, J., and Wang, J.: Size-resolved
560 airborne particulate oxalic and related secondary organic aerosol species in the
561 urban atmosphere of Chengdu, China, *Atmospheric Research*, 161, 134-142,
562 2015.

563 Cheng, C. L., Wang, G. H., Zhou, B. H., Meng, J. J., Li, J. J., Cao, J. J., and Xiao, S.:
564 Comparison of dicarboxylic acids and related compounds in aerosol samples
565 collected in Xi'an, China during haze and clean periods, *Atmospheric
566 Environment*, 81, 443-449, 10.1016/j.atmosenv.2013.09.013, 2013.

567 Cui, J. R., and Zhang, L. F.: Metallurgical recovery of metals from electronic waste: A
568 review, *J Hazard Mater*, 158, 228-256, 10.1016/j.jhazmat.2008.02.001, 2008.

569 Dall'Osto, M., Booth, M., Smith, W., Fisher, R., and Harrison, R. M.: A study of the
570 size distributions and the chemical characterization of airborne particles in the
571 vicinity of a large integrated steelworks, *Aerosol Science and Technology*, 42,
572 981-991, 2008.

573 Deguillaume, L., Leriche, M., Desboeufs, K., Mailhot, G., George, C., and
574 Chaumerliac, N.: Transition metals in atmospheric liquid phases: Sources,
575 reactivity, and sensitive parameters, *Chem Rev*, 105, 3388-3431,
576 10.1021/cr040649c, 2005.

577 Denkenberger, K. A., Moffet, R. C., Holecek, J. C., Rebotier, T. P., and Prather, K. A.:
578 Real-time, single-particle measurements of oligomers in aged ambient aerosol
579 particles, *Environmental Science & Technology*, 41, 5439-5446,
580 10.1021/es070329l, 2007.

581 Du, H., Kong, L., Cheng, T., Chen, J., Du, J., Li, L., Xia, X., Leng, C., and Huang, G.:
582 Insights into summertime haze pollution events over Shanghai based on online
583 water-soluble ionic composition of aerosols, *Atmospheric Environment*, 45,
584 5131-5137, 2011.

585 Dunford, H. B.: Oxidations of iron (II)/(III) by hydrogen peroxide: from aquo to
586 enzyme, *Coordination Chemistry Reviews*, 233, 311-318, 2002.

587 Ervens, B., Gligorovski, S., and Herrmann, H.: Temperature-dependent rate constants
588 for hydroxyl radical reactions with organic compounds in aqueous solutions,
589 *Phys Chem Chem Phys*, 5, 1811-1824, Doi 10.1039/B300072a, 2003.

590 Ervens, B., Feingold, G., Frost, G. J., and Kreidenweis, S. M.: A modeling study of
591 aqueous production of dicarboxylic acids: 1. Chemical pathways and speciated
592 organic mass production, *Journal of Geophysical Research-Atmospheres*, 109,
593 D15205, 10.1029/2003jd004387, 2004.

594 Ervens, B., and Volkamer, R.: Glyoxal processing by aerosol multiphase chemistry:
595 towards a kinetic modeling framework of secondary organic aerosol formation in
596 aqueous particles, *Atmospheric Chemistry and Physics*, 10, 8219-8244, DOI

597 10.5194/acp-10-8219-2010, 2010.

598 Ervens, B., Turpin, B. J., and Weber, R. J.: Secondary organic aerosol formation in
599 cloud droplets and aqueous particles (aqSOA): a review of laboratory, field and
600 model studies, *Atmospheric Chemistry and Physics*, 11, 11069-11102,
601 10.5194/acp-11-11069-2011, 2011.

602 Ervens, B., Sorooshian, A., Lim, Y. B., and Turpin, B. J.: Key parameters controlling
603 OH-initiated formation of secondary organic aerosol in the aqueous phase
604 (aqSOA), *Journal of Geophysical Research-Atmospheres*, 119, 3997-4016,
605 10.1002/2013JD021021, 2014.

606 Ervens, B.: Modeling the Processing of Aerosol and Trace Gases in Clouds and Fogs,
607 *Chem Rev*, 115, 4157-4198, 10.1021/cr5005887, 2015.

608 Fenton, H.: LXXIII.—Oxidation of tartaric acid in presence of iron, *Journal of the*
609 *Chemical Society, Transactions*, 65, 899-910, 1894.

610 Fountoukis, C., and Nenes, A.: ISORROPIA II: a computationally efficient
611 thermodynamic equilibrium model for K^+ - Ca^{2+} - Mg^{2+} - NH_4^+ - Na^+ - SO_4^{2-} - NO_3^- - Cl^- - H_2O aerosols, *Atmospheric Chemistry and Physics*, 7,
612 4639-4659, 2007.

614 Galloway, M. M., Chhabra, P. S., Chan, A. W. H., Surratt, J. D., Flagan, R. C.,
615 Seinfeld, J. H., and Keutsch, F. N.: Glyoxal uptake on ammonium sulphate seed
616 aerosol: reaction products and reversibility of uptake under dark and irradiated
617 conditions, *Atmospheric Chemistry and Physics*, 9, 3331-3345,
618 10.5194/acp-9-3331-2009, 2009.

619 Gligorovski, S., Streckowski, R., Barbati, S., and Vione, D.: Environmental
620 Implications of Hydroxyl Radicals ($\bullet OH$), *Chem Rev*, 115, 13051-13092,
621 10.1021/cr500310b, 2015.

622 Goldstein, A. H., and Galbally, I. E.: Known and unexplored organic constituents in
623 the earth's atmosphere, *Environmental Science & Technology*, 41, 1514-1521,
624 Doi 10.1021/Es072476p, 2007.

625 Hallquist, M., Wenger, J. C., Baltensperger, U., Rudich, Y., Simpson, D., Claeys, M.,
626 Dommen, J., Donahue, N. M., George, C., Goldstein, A. H., Hamilton, J. F.,
627 Herrmann, H., Hoffmann, T., Iinuma, Y., Jang, M., Jenkin, M. E., Jimenez, J. L.,
628 Kiendler-Scharr, A., Maenhaut, W., McFiggans, G., Mentel, T. F., Monod, A.,
629 Prévôt, A. S. H., Seinfeld, J. H., Surratt, J. D., Szmigielski, R., and Wildt, J.: The
630 formation, properties and impact of secondary organic aerosol: current and
631 emerging issues, *Atmospheric Chemistry and Physics*, 9, 5155-5236,
632 10.5194/acp-9-5155-2009, 2009.

633 Hanson, D. R., Burkholder, J. B., Howard, C. J., and Ravishankara, A.: Measurement
634 of hydroxyl and hydroperoxy radical uptake coefficients on water and sulfuric
635 acid surfaces, *The Journal of Physical Chemistry*, 96, 4979-4985, 1992.

636 Hatch, L. E., Creamean, J. M., Ault, A. P., Surratt, J. D., Chan, M. N., Seinfeld, J. H.,
637 Edgerton, E. S., Su, Y., and Prather, K. A.: Measurements of isoprene-derived
638 organosulfates in ambient aerosols by aerosol time-of-flight mass
639 spectrometry-Part 1: Single particle atmospheric observations in Atlanta,
640 *Environmental science & technology*, 45, 5105-5111, 2011.

641 Herrmann, H.: Kinetics of aqueous phase reactions relevant for atmospheric chemistry,
642 Chem Rev, 103, 4691-4716, Doi 10.1021/Cr020658q, 2003.

643 Herrmann, H., Tilgner, A., Barzaghi, P., Majdik, Z., Gligorovski, S., Poulain, L., and
644 Monod, A.: Towards a more detailed description of tropospheric aqueous phase
645 organic chemistry: CAPRAM 3.0, Atmospheric Environment, 39, 4351-4363,
646 10.1016/j.atmosenv.2005.02.016, 2005.

647 Herrmann, H., Schaefer, T., Tilgner, A., Styler, S. A., Weller, C., Teich, M., and Otto,
648 T.: Tropospheric Aqueous-Phase Chemistry: Kinetics, Mechanisms, and Its
649 Coupling to a Changing Gas Phase, Chem Rev, 115, 4259-4334,
650 10.1021/cr500447k, 2015.

651 Ho, K. F., Cao, J. J., Lee, S. C., Kawamura, K., Zhang, R. J., Chow, J. C., and Watson,
652 J. G.: Dicarboxylic acids, ketocarboxylic acids, and dicarbonyls in the urban
653 atmosphere of China, Journal of Geophysical Research-Atmospheres, 112,
654 D22S27, 10.1029/2006jd008011, 2007.

655 Ho, K. F., Lee, S. C., Ho, S. S. H., Kawamura, K., Tachibana, E., Cheng, Y., and Zhu,
656 T.: Dicarboxylic acids, ketocarboxylic acids, alpha-dicarbonyls, fatty acids, and
657 benzoic acid in urban aerosols collected during the 2006 Campaign of Air
658 Quality Research in Beijing (CAREBeijing-2006), Journal of Geophysical
659 Research-Atmospheres, 115, D19312, 10.1029/2009jd013304, 2010.

660 Ho, K. F., Ho, S. S. H., Lee, S. C., Kawamura, K., Zou, S. C., Cao, J. J., and Xu, H.
661 M.: Summer and winter variations of dicarboxylic acids, fatty acids and benzoic
662 acid in PM(2.5) in Pearl Delta River Region, China, Atmospheric Chemistry and
663 Physics, 11, 2197-2208, 10.5194/acp-11-2197-2011, 2011.

664 Huang, X.-F., and Yu, J. Z.: Is vehicle exhaust a significant primary source of oxalic
665 acid in ambient aerosols?, Geophysical Research Letters, 34, L02808,
666 10.1029/2006gl028457, 2007.

667 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q.,
668 Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C.,
669 Docherty, K. S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J.,
670 Smith, J. D., Wilson, K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J.,
671 Laaksonen, A., Raatikainen, T., Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala,
672 M., Tomlinson, J. M., Collins, D. R., Cubison, M. J., Dunlea, E. J., Huffman, J.
673 A., Onasch, T. B., Alfarra, M. R., Williams, P. I., Bower, K., Kondo, Y.,
674 Schneider, J., Drewnick, F., Borrmann, S., Weimer, S., Demerjian, K., Salcedo,
675 D., Cottrell, L., Griffin, R., Takami, A., Miyoshi, T., Hatakeyama, S., Shimojo,
676 A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J. R., Sueper, D., Jayne, J. T.,
677 Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood, E. C., Middlebrook, A.
678 M., Kolb, C. E., Baltensperger, U., and Worsnop, D. R.: Evolution of Organic
679 Aerosols in the Atmosphere, Science, 326, 1525-1529, 10.1126/science.1180353,
680 2009.

681 Kawamura, K., and Ikushima, K.: seasonal-changes in the distribution of
682 dicarboxylic-acids in the urban atmosphere, Environmental Science &
683 Technology, 27, 2227-2235, 10.1021/es00047a033, 1993.

684 Kawamura, K., Kobayashi, M., Tsubonuma, N., Mochida, M., Watanabe, T., and Lee,

685 M.: Organic and inorganic compositions of marine aerosols from East Asia:
686 Seasonal variations of water-soluble dicarboxylic acids, major ions, total carbon
687 and nitrogen, and stable C and N isotopic composition, *Geochemical*
688 *Investigations in Earth and Space Science: A Tribute to Issac R. Kaplan*, 9, edited
689 by: Hill, R. J. L. J. A. Z. B. M. J. C. G., and Eganhouse, R. G. M. P. K., 243-265
690 pp., 2004.

691 Kawamura, K., and Yasui, O.: Diurnal changes in the distribution of dicarboxylic
692 acids, ketocarboxylic acids and dicarbonyls in the urban Tokyo atmosphere,
693 *Atmospheric Environment*, 39, 1945-1960, 10.1016/j.atmosenv.2004.12.014,
694 2005.

695 Kawamura, K., Kasukabe, H., and Barrie, L. A.: Secondary formation of
696 water-soluble organic acids and alpha-dicarbonyls and their contributions to total
697 carbon and water-soluble organic carbon: Photochemical aging of organic
698 aerosols in the Arctic spring, *Journal of Geophysical Research-Atmospheres*, 115,
699 D21306, 10.1029/2010jd014299, 2010.

700 Kawamura, K., Tachibana, E., Okuzawa, K., Aggarwal, S. G., Kanaya, Y., and Wang,
701 Z. F.: High abundances of water-soluble dicarboxylic acids, ketocarboxylic acids
702 and alpha-dicarbonyls in the mountaintop aerosols over the North China Plain
703 during wheat burning season, *Atmospheric Chemistry and Physics*, 13,
704 8285-8302, 10.5194/acp-13-8285-2013, 2013.

705 Khezri, B., Mo, H., Yan, Z., Chong, S.-L., Heng, A. K., and Webster, R. D.:
706 Simultaneous online monitoring of inorganic compounds in aerosols and gases in
707 an industrialized area, *Atmospheric Environment*, 80, 352-360, 2013.

708 Kremer, M. L.: The Fenton reaction. Dependence of the rate on pH, *Journal of*
709 *Physical Chemistry A*, 107, 1734-1741, 10.1021/jp020654p, 2003.

710 Kundu, S., Kawamura, K., Andreae, T. W., Hoffer, A., and Andreae, M. O.: Molecular
711 distributions of dicarboxylic acids, ketocarboxylic acids and alpha-dicarbonyls in
712 biomass burning aerosols: implications for photochemical production and
713 degradation in smoke layers, *Atmospheric Chemistry and Physics*, 10, 2209-2225,
714 10.5194/acp-10-2209-2010, 2010.

715 Li, L., Huang, Z. X., Dong, J. G., Li, M., Gao, W., Nian, H. Q., Fu, Z., Zhang, G. H.,
716 Bi, X. H., Cheng, P., and Zhou, Z.: Real time bipolar time-of-flight mass
717 spectrometer for analyzing single aerosol particles, *Int J Mass Spectrom*, 303,
718 118-124, 10.1016/j.ijms.2011.01.017, 2011.

719 Li, L., Li, M., Huang, Z., Gao, W., Nian, H., Fu, Z., Gao, J., Chai, F., and Zhou, Z.:
720 Ambient particle characterization by single particle aerosol mass spectrometry in
721 an urban area of Beijing, *Atmospheric Environment*, 94, 323-331, 2014.

722 Lim, Y. B., Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Aqueous chemistry
723 and its role in secondary organic aerosol (SOA) formation, *Atmospheric*
724 *Chemistry and Physics*, 10, 10521-10539, 10.5194/acp-10-10521-2010, 2010.

725 Ma, Q. X., He, H., and Liu, C.: Hygroscopic properties of oxalic acid and
726 atmospherically relevant oxalates, *Atmospheric Environment*, 69, 281-288,
727 10.1016/j.atmosenv.2012.12.011, 2013.

728 Ma, S. X., Ren, K., Liu, X. W., Chen, L. G., Li, M., Li, X. Y., Yang, J., Huang, B.,

729 Zheng, M., and Xu, Z. C.: Production of hydroxyl radicals from Fe-containing
730 fine particles in Guangzhou, China, *Atmospheric Environment*, 123, 72-78,
731 10.1016/j.atmosenv.2015.10.057, 2015.

732 McNeill, V. F.: Aqueous organic chemistry in the atmosphere: Sources and chemical
733 processing of organic aerosols, *Environmental science & technology*, 49,
734 1237-1244, 2015.

735 Meng, J. J., Wang, G. H., Li, J. J., Cheng, C. L., and Cao, J. J.: Atmospheric oxalic
736 acid and related secondary organic aerosols in Qinghai Lake, a continental
737 background site in Tibet Plateau, *Atmospheric Environment*, 79, 582-589,
738 10.1016/j.atmosenv.2013.07.024, 2013.

739 Meng, J. J., Wang, G. H., Li, J. J., Cheng, C. L., Ren, Y. Q., Huang, Y., Cheng, Y. T.,
740 Cao, J. J., and Zhang, T.: Seasonal characteristics of oxalic acid and related SOA
741 in the free troposphere of Mt. Hua, central China: Implications for sources and
742 formation mechanisms, *Science Of The Total Environment*, 493, 1088-1097,
743 10.1016/j.scitotenv.2014.04.086, 2014.

744 Miyazaki, Y., Aggarwal, S. G., Singh, K., Gupta, P. K., and Kawamura, K.:
745 Dicarboxylic acids and water-soluble organic carbon in aerosols in New Delhi,
746 India, in winter: Characteristics and formation processes, *Journal of Geophysical
747 Research-Atmospheres*, 114, D19206, 10.1029/2009jd011790, 2009.

748 Moffet, R. C., de Foy, B., Molina, L. T., Molina, M. J., and Prather, K. A.:
749 Measurement of ambient aerosols in northern Mexico City by single particle
750 mass spectrometry, *Atmospheric Chemistry and Physics*, 8, 4499-4516,
751 10.5194/acp-8-4499-2008, 2008.

752 Myriokefalitakis, S., Tsigaridis, K., Mihalopoulos, N., Sciare, J., Nenes, A.,
753 Kawamura, K., Segers, A., and Kanakidou, M.: In-cloud oxalate formation in the
754 global troposphere: a 3-D modeling study, *Atmospheric Chemistry and Physics*,
755 11, 5761-5782, 10.5194/acp-11-5761-2011, 2011.

756 Nenes, A., Pandis, S. N., and Pilinis, C.: ISORROPIA: A new thermodynamic
757 equilibrium model for multiphase multicomponent inorganic aerosols, *Aquatic
758 geochemistry*, 4, 123-152, 1998.

759 Nenes, A., Pandis, S. N., and Pilinis, C.: Continued development and testing of a new
760 thermodynamic aerosol module for urban and regional air quality models,
761 *Atmospheric Environment*, 33, 1553-1560, 1999.

762 Noble, C. A., and Prather, K. A.: Real-time measurement of correlated size and
763 composition profiles of individual atmospheric aerosol particles, *Environmental
764 science & technology*, 30, 2667-2680, 1996.

765 Novakov, T., and Penner, J. E.: Large Contribution of Organic Aerosols to
766 Cloud-Condensation-Nuclei Concentrations, *Nature*, 365, 823-826, Doi
767 10.1038/365823a0, 1993.

768 Pavuluri, C. M., Kawamura, K., and Swaminathan, T.: Water-soluble organic carbon,
769 dicarboxylic acids, ketoacids, and alpha-dicarbonyls in the tropical Indian
770 aerosols, *Journal of Geophysical Research-Atmospheres*, 115, D11302,
771 10.1029/2009JD012661, 2010.

772 Poschl, U., and Shiraiwa, M.: Multiphase Chemistry at the Atmosphere-Biosphere

773 Interface Influencing Climate and Public Health in the Anthropocene, *Chem Rev*,
774 115, 4440-4475, 10.1021/cr500487s, 2015.

775 Prather, K. A., Nordmeyer, T., and Salt, K.: Real-time characterization of individual
776 aerosol particles using time-of-flight mass spectrometry, *Anal. Chem.*, 66,
777 1403-1407, 1994.

778 Pratt, K. A., Hatch, L. E., and Prather, K. A.: Seasonal Volatility Dependence of
779 Ambient Particle Phase Amines, *Environmental Science & Technology*, 43,
780 5276-5281, 10.1021/es803189n, 2009.

781 Prenni, A. J., De Mott, P. J., and Kreidenweis, S. M.: Water uptake of internally mixed
782 particles containing ammonium sulfate and dicarboxylic acids, *Atmospheric*
783 *Environment*, 37, 4243-4251, 10.1016/s1352-2310(03)00559-4, 2003.

784 Silva, P. J., and Prather, K. A.: Interpretation of mass spectra from organic compounds
785 in aerosol time-of-flight mass spectrometry, *Analytical Chemistry*, 72, 3553-3562,
786 2000.

787 Sorooshian, A., Varutbangkul, V., Brechtel, F. J., Ervens, B., Feingold, G., Bahreini,
788 R., Murphy, S. M., Holloway, J. S., Atlas, E. L., Buzorius, G., Jonsson, H.,
789 Flagan, R. C., and Seinfeld, J. H.: Oxalic acid in clear and cloudy atmospheres:
790 Analysis of data from International Consortium for Atmospheric Research on
791 Transport and Transformation 2004, *Journal of Geophysical*
792 *Research-Atmospheres*, 111, 10.1029/2005jd006880, 2006.

793 Sorooshian, A., Lu, M.-L., Brechtel, F. J., Jonsson, H., Feingold, G., Flagan, R. C.,
794 and Seinfeld, J. H.: On the source of organic acid aerosol layers above clouds,
795 *Environmental Science & Technology*, 41, 4647-4654, 10.1021/es0630442,
796 2007a.

797 Sorooshian, A., Ng, N. L., Chan, A. W. H., Feingold, G., Flagan, R. C., and Seinfeld, J.
798 H.: Particulate organic acids and overall water-soluble aerosol composition
799 measurements from the 2006 Gulf of Mexico Atmospheric Composition and
800 Climate Study (GoMACCS), *Journal of Geophysical Research-Atmospheres*,
801 112, 10.1029/2007jd008537, 2007b.

802 Sorooshian, A., Wang, Z., Coggon, M. M., Jonsson, H. H., and Ervens, B.:
803 Observations of Sharp Oxalate Reductions in Stratocumulus Clouds at Variable
804 Altitudes: Organic Acid and Metal Measurements During the 2011 E-PEACE
805 Campaign, *Environmental Science & Technology*, 47, 7747-7756,
806 10.1021/es4012383, 2013.

807 Stone, E. A., Hedman, C. J., Zhou, J. B., Mieritz, M., and Schauer, J. J.: Insights into
808 the nature of secondary organic aerosol in Mexico City during the MILAGRO
809 experiment 2006, *Atmospheric Environment*, 44, 312-319,
810 10.1016/j.atmosenv.2009.10.036, 2010.

811 Sullivan, R. C., and Prather, K. A.: Investigations of the diurnal cycle and mixing state
812 of oxalic acid in individual particles in Asian aerosol outflow, *Environmental*
813 *Science Technology*, 41, 8062-8069, 2007a.

814 Sullivan, R. C., and Prather, K. A.: Investigations of the diurnal cycle and mixing state
815 of oxalic acid in individual particles in Asian aerosol outflow, *Environmental*
816 *Science & Technology*, 41, 8062-8069, 10.1021/es071134g, 2007b.

817 Surratt, J. D., Kroll, J. H., Kleindienst, T. E., Edney, E. O., Claeys, M., Sorooshian, A.,
818 Ng, N. L., Offenberg, J. H., Lewandowski, M., Jaoui, M., Flagan, R. C., and
819 Seinfeld, J. H.: Evidence for organosulfates in secondary organic aerosol,
820 *Environmental Science & Technology*, 41, 517-527, 10.1021/Es062081q, 2007.

821 Surratt, J. D., Gomez-Gonzalez, Y., Chan, A. W. H., Vermeylen, R., Shahgholi, M.,
822 Kleindienst, T. E., Edney, E. O., Offenberg, J. H., Lewandowski, M., Jaoui, M.,
823 Maenhaut, W., Claeys, M., Flagan, R. C., and Seinfeld, J. H.: Organosulfate
824 formation in biogenic secondary organic aerosol, *Journal of Physical Chemistry*
825 *A*, 112, 8345-8378, Doi 10.1021/Jp802310p, 2008.

826 Tan, Y., Perri, M. J., Seitzinger, S. P., and Turpin, B. J.: Effects of Precursor
827 Concentration and Acidic Sulfate in Aqueous Glyoxal-OH Radical Oxidation and
828 Implications for Secondary Organic Aerosol, *Environmental Science &*
829 *Technology*, 43, 8105-8112, 10.1021/Es901742f, 2009.

830 ten Brink, H., Otjes, R., Jongejan, P., and Slanina, S.: An instrument for
831 semi-continuous monitoring of the size-distribution of nitrate, ammonium,
832 sulphate and chloride in aerosol, *Atmospheric Environment*, 41, 2768-2779,
833 2007.

834 Tong, H. J., Arangio, A. M., Lakey, P. S. J., Berkemeier, T., Liu, F. B., Kampf, C. J.,
835 Brune, W. H., Poschl, U., and Shiraiwa, M.: Hydroxyl radicals from secondary
836 organic aerosol decomposition in water, *Atmospheric Chemistry and Physics*, 16,
837 1761-1771, 10.5194/acp-16-1761-2016, 2016.

838 van Pinxteren, D., Neususs, C., and Herrmann, H.: On the abundance and source
839 contributions of dicarboxylic acids in size-resolved aerosol particles at
840 continental sites in central Europe, *Atmospheric Chemistry and Physics*, 14,
841 3913-3928, 10.5194/acp-14-3913-2014, 2014.

842 Wang, G., Cheng, C., Meng, J., Huang, Y., Li, J., and Ren, Y.: Field observation on
843 secondary organic aerosols during Asian dust storm periods: Formation
844 mechanism of oxalic acid and related compounds on dust surface, *Atmospheric*
845 *Environment*, 113, 169-176, 2015.

846 Wang, G., Zhang, R., Gomez, M. E., Yang, L., Levy Zamora, M., Hu, M., Lin, Y.,
847 Peng, J., Guo, S., Meng, J., Li, J., Cheng, C., Hu, T., Ren, Y., Wang, Y., Gao, J.,
848 Cao, J., An, Z., Zhou, W., Li, G., Wang, J., Tian, P., Marrero-Ortiz, W., Secret, J.,
849 Du, Z., Zheng, J., Shang, D., Zeng, L., Shao, M., Wang, W., Huang, Y., Wang, Y.,
850 Zhu, Y., Li, Y., Hu, J., Pan, B., Cai, L., Cheng, Y., Ji, Y., Zhang, F., Rosenfeld, D.,
851 Liss, P. S., Duce, R. A., Kolb, C. E., and Molina, M. J.: Persistent sulfate
852 formation from London Fog to Chinese haze, *Proceedings of the National*
853 *Academy of Sciences*, 113, 13630-13635, 10.1073/pnas.1616540113, 2016.

854 Wang, G. H., Kawamura, K., Cheng, C. L., Li, J. J., Cao, J. J., Zhang, R. J., Zhang, T.,
855 Liu, S. X., and Zhao, Z. Z.: Molecular Distribution and Stable Carbon Isotopic
856 Composition of Dicarboxylic Acids, Ketocarboxylic Acids, and
857 alpha-Dicarbonyls in Size-Resolved Atmospheric Particles From Xi'an City,
858 China, *Environmental Science & Technology*, 46, 4783-4791,
859 10.1021/es204322c, 2012.

860 Wang, J. Y., Wang, G. H., Gao, J., Wang, H., Ren, Y. Q., Li, J. J., Wu, C., Zhang, L.,

861 Wang, S. L., and Chai, F. H.: Concentrations and stable carbon isotope
862 compositions of oxalic acid and related SOA in Beijing before, during and after
863 the 2014 APEC, *Atmospheric Chemistry and Physics* 17, 981-992, 2017.

864 Weller, C., Tilgner, A., Brauer, P., and Herrmann, H.: Modeling the Impact of
865 Iron-Carboxylate Photochemistry on Radical Budget and Carboxylate
866 Degradation in Cloud Droplets and Particles, *Environmental Science &
867 Technology*, 48, 5652-5659, 10.1021/es4056643, 2014.

868 Wonaschuetz, A., Sorooshian, A., Ervens, B., Chuang, P. Y., Feingold, G., Murphy, S.
869 M., de Gouw, J., Warneke, C., and Jonsson, H. H.: Aerosol and gas
870 re-distribution by shallow cumulus clouds: An investigation using airborne
871 measurements, *Journal of Geophysical Research-Atmospheres*, 117,
872 10.1029/2012jd018089, 2012.

873 Xue, J., Lau, A. K. H., and Yu, J. Z.: A study of acidity on PM_{2.5} in Hong Kong using
874 online ionic chemical composition measurements, *Atmospheric Environment*, 45,
875 7081-7088, 10.1016/j.atmosenv.2011.09.040, 2011.

876 Yang, F., Chen, H., Wang, X., Yang, X., Du, J., and Chen, J.: Single particle mass
877 spectrometry of oxalic acid in ambient aerosols in Shanghai: Mixing state and
878 formation mechanism, *Atmospheric Environment*, 43, 3876-3882, 2009.

879 Yao, X. H., Fang, M., and Chan, C. K.: Size distributions and formation of
880 dicarboxylic acids in atmospheric particles, *Atmospheric Environment*, 36,
881 2099-2107, 2002.

882 Yao, X. H., Lau, A. P. S., Fang, M., Chan, C. K., and Hu, M.: Size distributions and
883 formation of ionic species in atmospheric particulate pollutants in Beijing, China:
884 2 - dicarboxylic acids, *Atmospheric Environment*, 37, 3001-3007,
885 10.1016/s1352-2310(03)00256-5, 2003.

886 Yu, J. Z., Huang, X. F., Xu, J. H., and Hu, M.: When aerosol sulfate goes up, so does
887 oxalate: Implication for the formation mechanisms of oxalate, *Environmental
888 Science & Technology*, 39, 128-133, 10.1021/Es049559f, 2005.

889 Yu, L., Smith, J., Laskin, A., Anastasio, C., Laskin, J., and Zhang, Q.: Chemical
890 characterization of SOA formed from aqueous-phase reactions of phenols with
891 the triplet excited state of carbonyl and hydroxyl radical, *Atmospheric Chemistry
892 and Physics*, 14, 13801-13816, 10.5194/acp-14-13801-2014, 2014.

893 Zauscher, M. D., Wang, Y., Moore, M. J. K., Gaston, C. J., and Prather, K. A.: Air
894 Quality Impact and Physicochemical Aging of Biomass Burning Aerosols during
895 the 2007 San Diego Wildfires, *Environmental Science & Technology*, 47,
896 7633-7643, 10.1021/es4004137, 2013.

897 Zhang, G., Bi, X., Li, L., Chan, L. Y., Li, M., Wang, X., Sheng, G., Fu, J., and Zhou,
898 Z.: Mixing state of individual submicron carbon-containing particles during
899 spring and fall seasons in urban Guangzhou, China: a case study, *Atmospheric
900 Chemistry and Physics*, 13, 4723-4735, 2013.

901 Zhang, G., Bi, X., He, J., Chen, D., Chan, L. Y., Xie, G., Wang, X., Sheng, G., Fu, J.,
902 and Zhou, Z.: Variation of secondary coatings associated with elemental carbon
903 by single particle analysis, *Atmospheric Environment*, 92, 162-170, 2014.

904 Zhang, R., Wang, G., Guo, S., Zamora, M. L., Ying, Q., Lin, Y., Wang, W., Hu, M.,

905 and Wang, Y.: Formation of urban fine particulate matter, *Chem Rev*, 115,
906 3803-3855, 2015.

907 Zhou, Y., Huang, X. H., Bian, Q., Griffith, S. M., Louie, P. K., and Yu, J. Z.: Sources
908 and atmospheric processes impacting oxalate at a suburban coastal site in Hong
909 Kong: Insights inferred from 1 year hourly measurements, *Journal of*
910 *Geophysical Research: Atmospheres*, 120, 9772-9788, 2015.

911

912

913

914

915

916

917

918

919

920

921

922

923

924

925

926

927

928

929

930

931

932

933

934

935

936

937

938

939

940

941

942

943

944

945

946

947

948

949 **Tables and Figures**

950

951 **Table list:**

952

953 Table 1. Summary of major groups of oxalic acid-containing particles in summer and
954 winter in PRD, China.

955

956 Table 2. The abundance of major particle types in total oxalic acid-containing
957 particles during the episode in winter (2/8/2015).

958

959 **Figure caption:**

960

961 Figure 1. Temporal variations of total detected particles and oxalic acid containing
962 particles during whole sampling periods in Heshan, China: (a) hourly variations of
963 PM_{2.5} mass concentration, total detected particle counts, oxalic acid containing
964 particles, ratio of oxalic acid-containing/total particles and major types of oxalic acid
965 containing particles; (b) variation patterns of relative abundance of major types of
966 oxalic acid containing particles.

967

968 Figure 2. The averaged positive and negative ion mass spectra of oxalic acid
969 containing particles is investigated in summer and winter: (a) summer positive, (b)
970 summer negative, (c) winter positive, (d) winter negative. The color bars represent
971 each peak area corresponding to specific fraction in individual particles.

972

973 Figure 3. (a) Mixing state of oxalic acid with sulfate, nitrate and ammonium in oxalic
974 acid-containing particles; (b) Linear correlation between NH₄⁺-containing oxalic acid
975 particles and the total oxalic acid particles in summer; (c) Linear correlation between
976 NH₄⁺-containing oxalic acid particles and the total oxalic acid particles in winter.
977 Abbreviations: C₂-NH₄⁺ represents the NH₄⁺-containing oxalic acid particles, and
978 same expressions for C₂-SO₄²⁻ and C₂-NO₃⁻.

979

980 Figure 4. Unscaled size-resolved number distributions of major types of oxalic acid
981 particles in summer and winter.

982

983 Figure 5. Temporal variations of O₃ concentrations, oxalic acid particles, malonic acid
984 particles and heavy metal type of oxalic acid particles during the entire sampling
985 period in Heshan, China.

986

987 Figure 6. The averaged digitized positive and negative ion mass spectra of heavy
988 metal type of oxalic acid-containing particles in summer.

989

990 Figure 7. The diurnal variations of O₃ concentration, oxalic acid particles, HM group
991 particles and in-situ pH (pH_{is}) from July 28 to August 1 in 2014.

992

993 Figure 8. The temporal variations of peak area of nitrate, sulfate and oxalic acid, and
994 the relative acidity ratio (R_{ra}) in carbonaceous type oxalic acid particles in winter.

995

996 Figure 9. The comprehensive study of oxalic acid particles increase on Feb 8, 2015: (a)
997 The digitized positive and negative ion mass spectrum of oxalic acid particles during
998 the episode; (b) Linear regression between total oxalic acid particles and
999 organosulfate-containing oxalic acid particles (m/z -155).

1000

1001

1002

1003

1004

1005

1006

1007

1008

1009

1010

Table 1. Summary of major groups of oxalic acid-containing particles in summer and winter in PRD, China.

Particle type	Summer(7/18-8/1, 2014)		Winter(1/27-2/8, 2015)	
	Count	Percentage, %	Count	Percentage, %
EC	1473	11.2	3161	15.4
ECOC	41	0.3	2233	10.9
OC	473	3.6	1922	9.4
BB	1702	13.0	4953	24.2
HM	4104	31.3	3124	15.2
Sec	2511	19.2	2192	10.7
NaK	303	2.3	17	0.1
Dust	1139	8.7	1888	9.2

Abbreviations of major particle types: elemental carbon (EC), elemental and organic carbon (ECOC), organic carbon (OC), biomass burning (BB), heavy metal (HM), secondary (Sec), sodium and potassium (NaK) and dust (Dust).

1011

1012

1013

1014

1015

Table 2. The abundance of major particle types in total oxalic acid-containing particles during the episode in winter (2/8/2015).

	EC	ECOC	OC	BB	Sec	HM	Dust	other
Count	1250	604	326	1320	856	377	814	132
Percentage, %	22.0	10.6	5.7	23.2	15.1	6.6	14.3	2.3

1016

1017

1018

1019

1020

1021

1022

1023

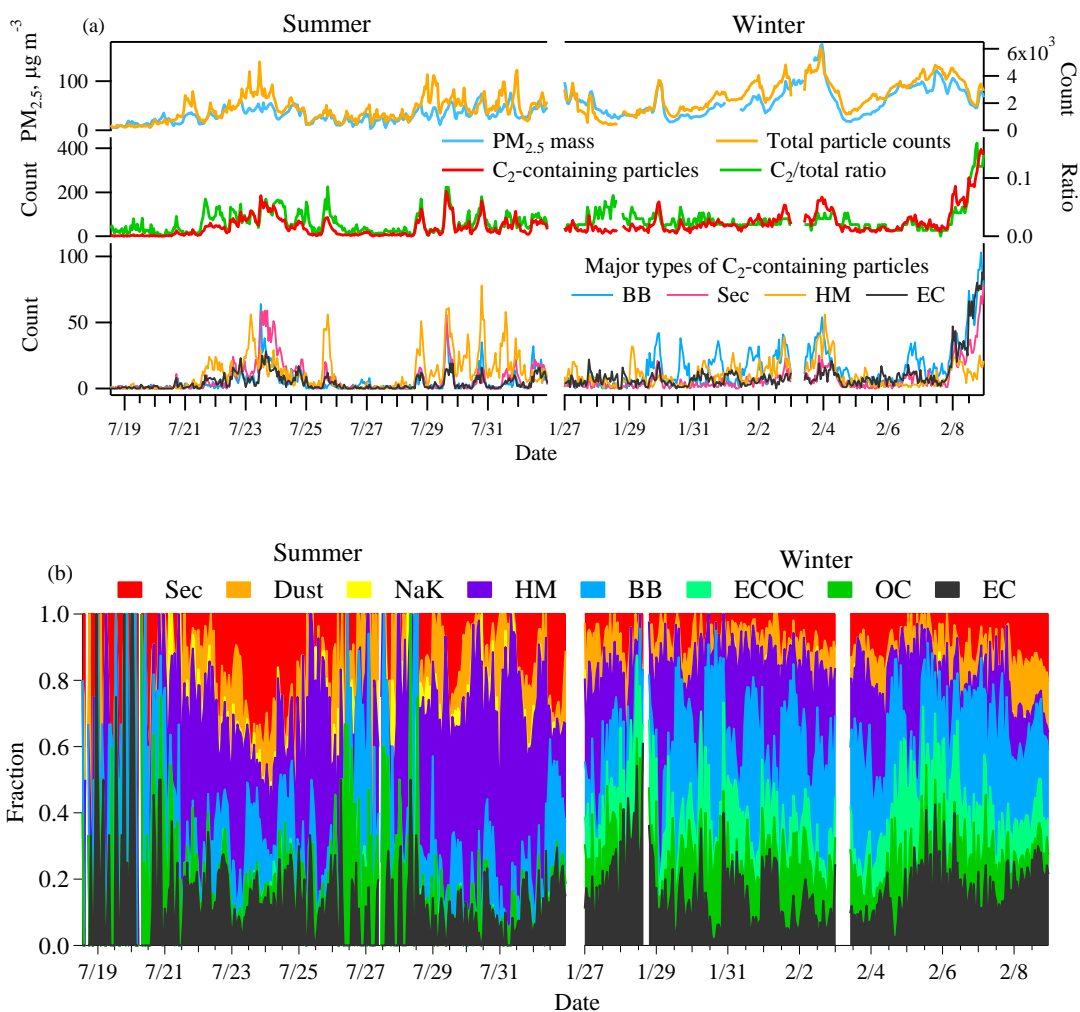
1024

1025

1026

1027

1028

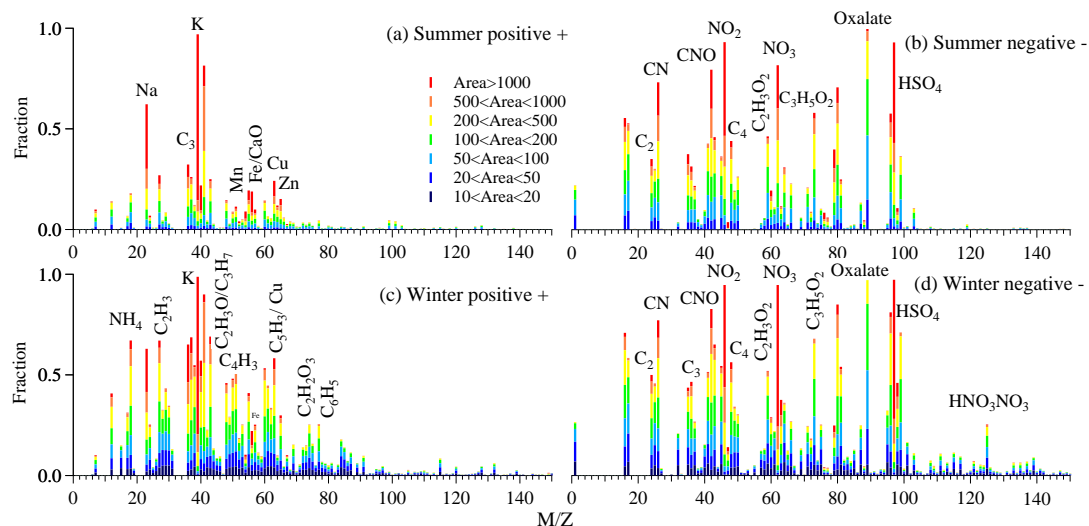


1029
1030

1031

1032 Figure 1. Temporal variations of total detected particles and oxalic acid containing
 1033 particles during whole sampling periods in Heshan, China: (a) hourly variations of
 1034 PM_{2.5} mass concentration, total detected particle counts, oxalic acid containing
 1035 particles, ratio of oxalic acid-containing/total particles and major types of oxalic acid
 1036 containing particles; (b) variation patterns of relative abundance of major types of
 1037 oxalic acid containing particles. Abbreviations of major particle types: elemental
 1038 carbon (EC), organic carbon (OC), elemental and organic carbon (ECOC), biomass
 1039 burning (BB), heavy metal (HM), secondary (Sec), sodium and potassium (NaK) and
 1040 dust.

1041
1042
1043
1044
1045



1046

1047 Figure 2. The averaged positive and negative ion mass spectra of oxalic acid
 1048 containing particles is investigated in summer and winter: (a) summer positive, (b)
 1049 summer negative, (c) winter positive, (d) winter negative. The color bars represent
 1050 each peak area corresponding to specific fraction in individual particles.

1051

1052

1053

1054

1055

1056

1057

1058

1059

1060

1061

1062

1063

1064

1065

1066

1067

1068

1069

1070

1071

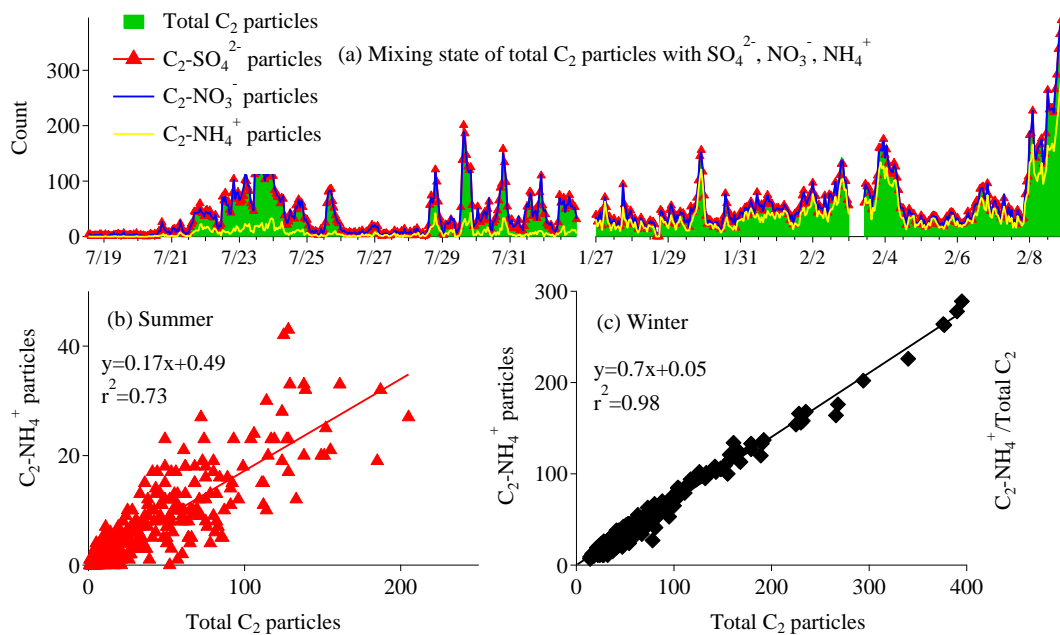
1072

1073

1074

1075

1076
1077



1078

1079 Figure 3. (a) Mixing state of oxalic acid with sulfate, nitrate and ammonium in oxalic
1080 acid-containing particles; (b) Linear correlation between NH₄⁺-containing oxalic acid
1081 particles and the total oxalic acid particles in summer; (c) Linear correlation between
1082 NH₄⁺-containing oxalic acid particles and the total oxalic acid particles in winter.
1083 Abbreviations: C₂-NH₄⁺ represents the NH₄⁺-containing oxalic acid particles, and
1084 same expressions for C₂-SO₄²⁻ and C₂-NO₃⁻.

1085

1086

1087

1088

1089

1090

1091

1092

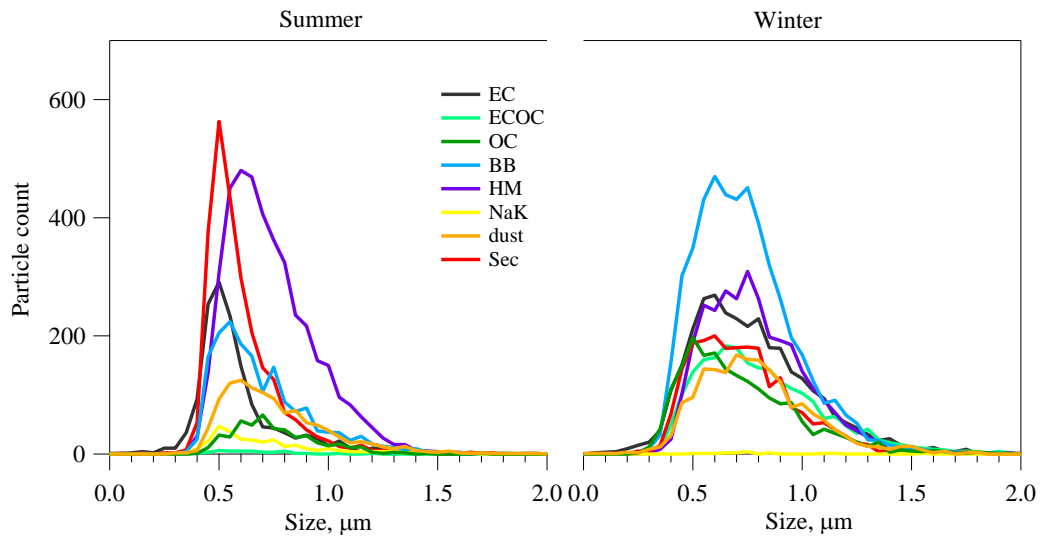
1093

1094

1095

1096

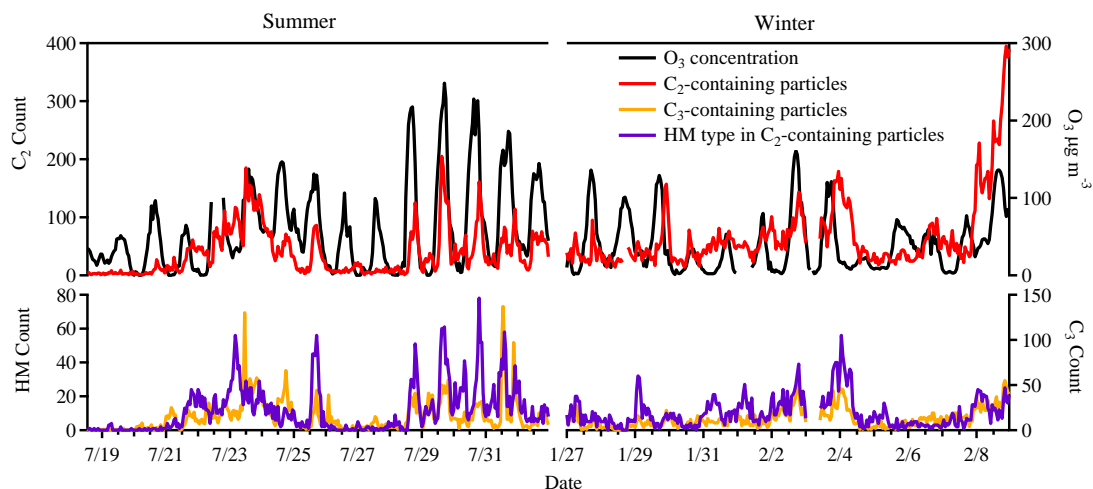
1097
1098
1099



1100
1101
1102
1103
1104
1105
1106
1107
1108
1109
1110
1111
1112
1113
1114
1115
1116
1117
1118
1119
1120
1121
1122
1123
1124

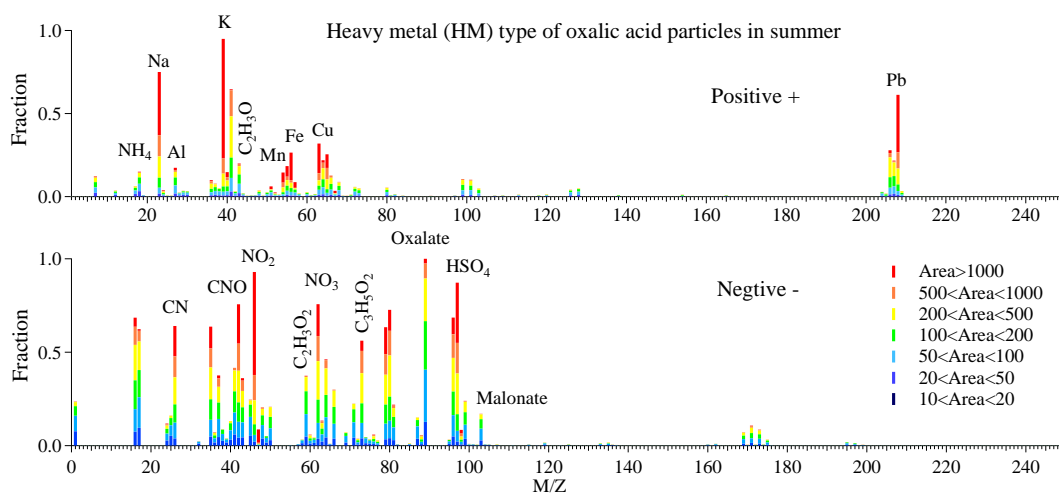
Figure 4. Unscaled size-resolved number distributions of major types of oxalic acid particles in summer and winter. Abbreviations of major particle types: elemental carbon (EC), organic carbon (OC), elemental and organic carbon (ECOc), biomass burning (BB), heavy metal (HM), secondary (Sec), sodium and potassium (NaK) and dust.

1125
1126



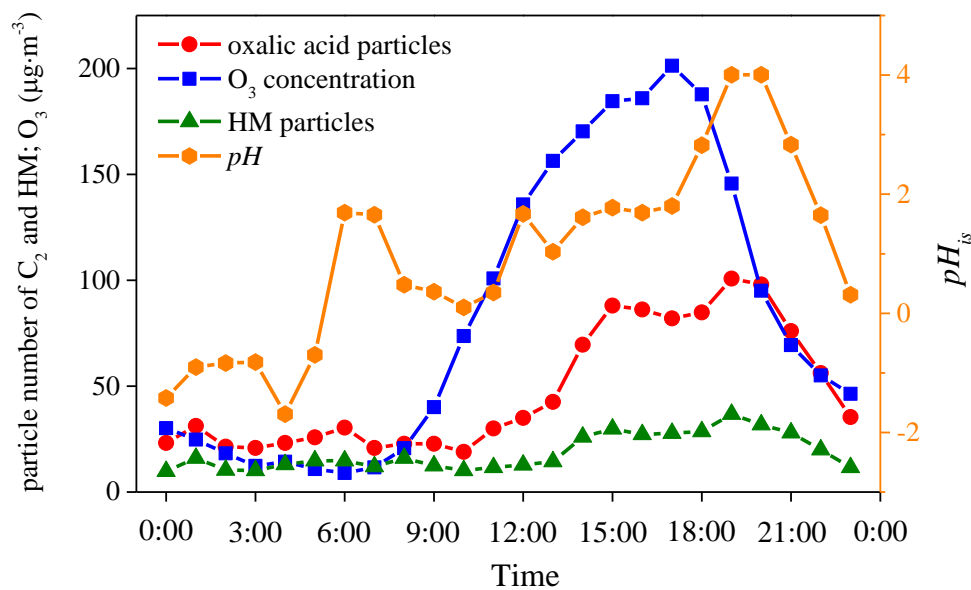
1127
1128
1129
1130
1131
1132
1133
1134

Figure 5. Temporal variations of O₃ concentrations, oxalic acid particles, malonic acid particles and heavy metal type of oxalic acid particles during the entire sampling period in Heshan, China.



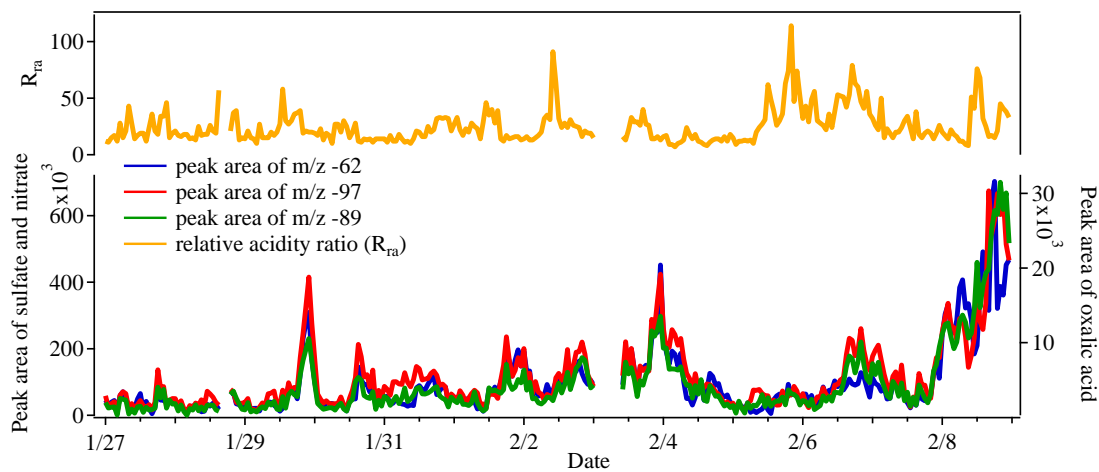
1135
1136
1137
1138
1139
1140

Figure 6. The averaged digitized positive and negative ion mass spectra of heavy metal type of oxalic acid-containing particles in summer.



1141
 1142
 1143
 1144
 1145
 1146
 1147

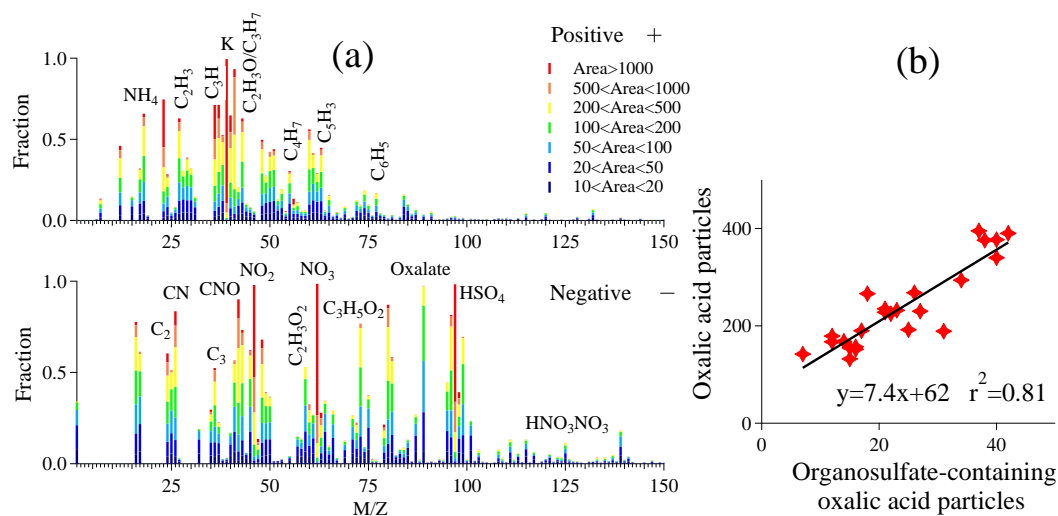
Figure 7. The diurnal variations of O₃ concentration, oxalic acid particles, HM group particles and in-situ pH (pH_{is}) from July 28 to August 1 in 2014.



1148
 1149
 1150
 1151
 1152
 1153
 1154

Figure 8. The temporal variations of peak area of nitrate, sulfate and oxalic acid, and the relative acidity ratio (R_{ra}) in carbonaceous type oxalic acid particles in winter.

1155



1156

1157 Figure 9. The comprehensive study of oxalic acid particles increase on Feb 8, 2015: (a)

1158 The digitized positive and negative ion mass spectrum of oxalic acid particles during

1159 the episode; (b) Linear regression between total oxalic acid particles and

1160 organosulfate-containing oxalic acid particles (m/z -155).

1161

1162