

1 Wintertime aerosol chemistry and haze evolution in an extremely 2 polluted city of North China Plain: significant contribution from coal 3 and biomass combustion

4 Haiyan Li^{1,2}, Qi Zhang², Qiang Zhang^{3,4}, Chunrong Chen³, Litao Wang⁵, Zhe Wei⁵, Shan Zhou², Caroline
5 Parworth², Bo Zheng¹, Francesco Canonaco⁶, André S. H. Prévôt⁶, Ping Chen⁷, Hongliang Zhang⁷,
6 Timothy J. Wallington⁸, Kebin He^{1,4,9}

7 ¹ State Key Joint Laboratory of Environment Simulation and Pollution Control, School of Environment, Tsinghua University,
8 Beijing 100084, China

9 ² Department of Environmental Toxicology, University of California, Davis, CA 95616, USA

10 ³ Ministry of Education Key Laboratory for Earth System Modeling, Center for Earth System Science, Tsinghua University,
11 Beijing 100084, China

12 ⁴ Collaborative Innovation Center for Regional Environmental Quality, Beijing 100084, China

13 ⁵ Department of Environmental Engineering, Hebei University of Engineering, Handan, Hebei 056038, China

14 ⁶ Laboratory of Atmospheric Chemistry, Paul Scherrer Institute, 5232 PSI Villigen, Switzerland

15 ⁷ Handix LLC, Boulder, CO 8031, USA

16 ⁸ Research & Advanced Engineering, Ford Motor Company, Dearborn, Michigan 28121, USA

17 ⁹ State Environmental Protection Key Laboratory of Sources and Control of Air Pollution Complex, Tsinghua University,
18 Beijing 100084, China

19 *Correspondence to:* Qiang Zhang (qiangzhang@tsinghua.edu.cn), Kebin He (hekb@tsinghua.edu.cn)

20 **Abstract.** The North China Plain (NCP) frequently experiences heavy haze pollution, particularly during wintertime. In 2015-
21 2016 winter, the NCP region suffered several extremely severe haze episodes with air pollution red alerts issued in many cities.
22 We have investigated the sources and aerosol evolution processes of the severe pollution episodes in Handan, a typical
23 industrialized city in the NCP region, using real-time measurements from an intensive field campaign during the winter of
24 2015-2016. The average ($\pm 1\sigma$) concentration of submicron aerosol (PM₁) during December 3, 2015 – February 5, 2016 was
25 187.6 (± 137.5) $\mu\text{g m}^{-3}$, with the hourly maximum reaching 700.8 $\mu\text{g m}^{-3}$. Organic was the most abundant component, on
26 average accounting for 45% of total PM₁ mass, followed by sulfate (15%), nitrate (14%), ammonium (12%), chloride (9%)
27 and BC (5%). Positive matrix factorization (PMF) with multi-linear engine (ME-2) identified four major organic aerosol (OA)
28 sources, including traffic emissions represented by a hydrocarbon-like OA (HOA, 7% of total OA), industrial and residential
29 burning of coal represented by a coal combustion OA (CCOA, 29% of total OA), open and domestic combustion of wood and
30 crop residuals represented by a biomass burning OA (BBOA, 25% of total OA), and formation of secondary OA (SOA) in the
31 atmosphere represented by an oxygenated OA (OOA, 39% of total OA). Emissions of primary OA (POA), which together
32 accounted for 61% of total OA and 27% of PM₁, are a major cause of air pollution during the winter. Our analysis further
33 uncovered that, primary emissions from coal combustion and biomass burning together with secondary formation of sulfate
34 (mainly from SO₂ emitted by coal combustion) are important driving factors for haze evolution. However, the bulk composition
35 of PM₁ showed comparatively small variations between less polluted periods (daily PM_{2.5} $\leq 75 \mu\text{g m}^{-3}$) and severely polluted

36 periods (daily $PM_{2.5} > 75 \mu g m^{-3}$), indicating relatively synchronous increases of all aerosol species during haze formation.
37 The case study of a severe haze episode, which lasted 8 days starting with a steady build-up of aerosol pollution followed by
38 a persistently high level of PM_1 ($326.7 - 700.8 \mu g m^{-3}$), revealed the significant influence of stagnant meteorological conditions
39 which exacerbate air pollution in the Handan region. The haze episode ended with a shift of wind which brought in cleaner air
40 masses from the northwest of Handan and gradually reduced PM_1 concentration to $< 50 \mu g m^{-3}$ after 12 hours. Aqueous-phase
41 reactions under higher relative humidity (RH) were found to significantly promote the production of secondary inorganic
42 species (especially sulfate), but showed little influence on SOA.

43 **1 Introduction**

44 Atmospheric particles are a complex mixture of species emitted directly to the atmosphere or formed via gas-to-particle
45 conversions. Aerosols can reduce visibility, adversely affect human health (Pope and Dockery, 2006), and influence climate
46 change directly by absorbing and reflecting solar radiation and indirectly by modifying cloud formation and properties (Pöschl,
47 2005; Seinfeld and Pandis, 2012), all of which are intrinsically linked to the chemical composition of aerosols. Therefore, it is
48 crucial to gain a quantitative understanding of aerosol composition and evolution processes for accurately assessing the
49 environmental effects of aerosols.

50 With the rapid economy growth and urbanization in North China Plain (NCP), air pollution in this region has become a
51 severe problem and a source of concern. Hebei Province, located in the NCP region, is known for persistent air quality problems
52 and extreme haze pollution events. According to the Ministry of Environmental Protection (MEP) of China, 7 out of the top
53 10 polluted cities in China in 2015 were located in Hebei province. During the extremely severe haze event that occurred in
54 the winter of 2015-2016 in the NCP region, the hourly peak $PM_{2.5}$ concentration in southern Hebei exceeded $1000 \mu g/m^3$. It is
55 well known that the severe air pollution in the NCP region was caused by large anthropogenic emissions and unfavorable
56 meteorological conditions. Emissions of primary $PM_{2.5}$, sulfur dioxide (SO_2), and nitrogen oxides (NO_x) from Hebei in 2015
57 are estimated to account for 8 %, 6 %, and 7 % of China's national total emissions respectively (<http://meicmodel.org/>), with
58 large contributions from coal and biomass combustion.

59 Large anthropogenic emissions in the NCP region have degraded regional air quality significantly. Extensive studies have
60 been conducted to explore the sources and evolution of haze episodes in Beijing, especially with the wide application of
61 Aerodyne Aerosol Mass Spectrometer (AMS)/Aerosol Chemical Speciation Monitor (ACSM) for online measurement of
62 aerosol chemical composition (Takegawa et al., 2009; Sun et al., 2010; Sun et al., 2012, 2013a, 2013b, 2014, 2015, 2016a,
63 2016b; Zhang et al., 2014; Hu et al., 2016). These studies have noted that regional air transport from the south or east
64 surrounding regions, unfavorable synoptic conditions, and heterogeneous secondary reactions associated with high RH
65 initiated the rapid formation and persistent evolution of haze episodes in Beijing. During a record-breaking haze episode in
66 wintertime in Beijing, Sun et al. (2014) estimated that regional transport contributed up to 66% to the steep rise of air pollutants
67 in Beijing. New particle formation and growth also plays an important role in haze formation. By examining in detail the haze

68 events under typical fall conditions in Beijing, Guo et al. (2014) indicated that nucleation consistently preceded a polluted
69 period with high number concentrations and the development of the episode involved efficient and sustained growth from the
70 nucleation-mode particles over multiple days. In addition, organic aerosol (OA) was found to be a major component of aerosol
71 particles, accounting for more than one-third of total PM₁ mass. The primary OA (POA) from traffic, cooking, biomass burning,
72 coal combustion, etc., and secondary OA (SOA) have been distinguished and quantified mainly using positive matrix
73 factorization (PMF; Paatero and Tapper, 1994). Recently, a novel PMF procedure, with the multi-linear engine (ME-2)
74 algorithm, was developed to apportion the OA sources in Beijing and Xi'an, allowing for a more objective selection of source
75 apportionment solution (Elser et al., 2016). However, our knowledge of the sources and aerosol evolution processes for the
76 whole region still remains incomplete and is especially limited for areas outside of Beijing. For other cities in the NCP region,
77 such as Hebei province, only a limited number of aerosol studies have been conducted using offline filter-based measurement
78 techniques (Zhao et al., 2013; Wei et al., 2014). Due to low time resolution varying from one day to several days, these studies
79 provided relatively limited information on aerosol emission sources and formation processes, thus it remains unclear how the
80 rapid haze evolution happens and what the driving sources are for the air pollution problems in Hebei. Therefore, it is crucial
81 to conduct research in the areas outside of Beijing, especially many provinces subjected to high anthropogenic emissions,
82 which may provide critical information to help air pollution policy making to be more direct and efficient.

83 To fill this knowledge gap, an intensive field campaign with multiple state-of-the-art research instruments was conducted
84 in Handan, a major city in southern Hebei, during the winter of 2015-2016. Handan is located in the intersectional area of four
85 provinces, Hebei, Shanxi, Henan and Shandong, all of which are heavily urbanized and industrialized (Fig. 1a). Handan itself
86 is also well known for heavy industrial production of steel, iron and cement, which result in high local emissions of air
87 pollutants. According to the routine monitoring of the China National Environmental Monitoring Center (CNEMC) from 2013
88 to 2015, Handan is always listed as one of the top 10 polluted cities in China. Hence, this location and its specific conditions
89 allow for a detailed exploration of aerosol chemistry and haze evolution processes under high anthropogenic emissions.

90 Here, we provide both overview and evolution cycle analyses of aerosol characteristics using aerosol data acquired with
91 an ACSM and collocated measurements of black carbon (BC), meteorological conditions and gas phase species. The sources
92 of OA are investigated in detail using PMF solved with the ME-2 algorithm (Paatero, 1999). Comparison of species diurnal
93 cycles between weekdays and weekends, polluted and non-polluted days, and the variation of aerosol characteristics with
94 increasing PM₁ concentration, provide insights into the driving factors for haze evolution. We also examine the impacts of
95 meteorological conditions based on an intense evolution case of submicron aerosol.

96 **2 Experimental methods**

97 **2.1 Sampling site and instrumentation**

98 In situ measurements were conducted at Hebei University of Engineering (36.57° N, 114.50° E) in Handan from

99 December 3, 2015 to February 5, 2016. Our sampling site is situated in the southeast edge of urban Handan, on the roof of a
100 four-story building (~12 m high), surrounded by the school and residential area, ~300 m north of the South Ring Road, and
101 ~400 m northeast of Handa Highway (S313). The ambient temperature varied from -12.7 to 14.4 °C, with an average of 1.8
102 °C. The prevailing wind came from the northeast and southwest, characterized by low wind speeds (Fig. 1b).

103 The mass concentrations of non-refractory submicron aerosol (NR-PM₁), including organics, sulfate, nitrate, ammonium,
104 and chloride, were measured in situ using an Aerodyne ACSM. The detailed description of this instrument can be found in Ng
105 et al. (2011a). In brief, ambient air was sampled through a PM_{2.5} cyclone to remove coarse particles with diameters exceeding
106 2.5 μm and then traversed a 2-m-long, ½-inch (outer diameter) stainless steel tube at a flow rate of 3 L min⁻¹ using an external
107 pump. A Nafion dryer was installed before the ACSM to dry aerosol samples and maintain the RH below 30%. Subsequently,
108 only a subset of the flow at ~85 cc min⁻¹ was sampled through a 100 μm critical orifice, focusing aerosol particles between 40
109 nm and 1 μm into the vacuum chamber via an aerodynamic lens. In our study, the ACSM mass spectrometer was operated at
110 a scanning speed of 200 ms per amu from 10 to 150 amu. By automatically switching 14 cycles between filter mode and
111 sample mode, the time resolution for the ACSM data in this study was approximately 15 minutes.

112 Because of the limit of the vaporizer temperature (~600 °C), the ACSM could not measure refractory species such as BC.
113 Thus a multi-angle absorption photometer (MAAP, Thermo Scientific model 5012) was deployed for real-time measurement
114 of BC concentration. The MAAP was operated at an incident light wavelength of 670 nm, with a PM₁ cyclone and a drying
115 system incorporated in front of the sampling line (Petzold and Schönlinner, 2004; Petzold et al., 2005). Online PM_{2.5} mass
116 concentration was measured simultaneously using a heated Tapered Element Oscillating Microbalance (TEOM series 1400a,
117 Thermo Scientific). Other collocated instruments included a suite of commercial gas analyzers (Thermo Scientific) to monitor
118 the variations of gaseous species (i.e., CO, O₃, NO, NO₂, NO_x, and SO₂). Meteorological parameters, i.e. temperature, RH,
119 pressure, wind speed (WS) and wind direction (WD), were obtained by a Lufft WS500-UMB Smart Weather Sensor. The data
120 reported in this paper are in Beijing Time (BJT: UTC+8).

121 **2.2 ACSM Data analysis**

122 The mass concentrations of non-refractory aerosol species and the spectral matrices of OA were processed using ACSM
123 standard data analysis software (v 1.5.3.5) within Igor Pro version 6.37. The detailed procedures have been described in Ng et
124 al. (2011a). The default relative ionization efficiency (RIE) values were used for organics (1.4), sulfate (1.19), nitrate (1.1),
125 and chloride (1.3), whereas the RIE of ammonium (6.28) was directly determined via analyzing pure NH₄NO₃ particles. To
126 account for the incomplete detection of aerosol species, a default collection efficiency (CE) value of 0.5 was applied to the
127 entire data set as aerosol particles were dried before ACSM sampling and the ammonium nitrate fraction was always lower
128 than 0.4 during the whole period. Although previous studies have shown that aerosol particles may be slightly acidic during
129 wintertime in the NCP region, particle acidity was not high enough to affect CE values substantially (Sun et al., 2016a). As
130 shown in Fig. S1 in the supplementary information, the mass concentrations of PM₁ (= NR-PM₁ + BC) correlated tightly with
131 total PM_{2.5} mass loadings measured by TEOM (slope=0.88, r=0.93). Compared to the results reported previously in this area

132 (Sun et al., 2013a; Sun et al., 2014; Zhang et al., 2014; Sun et al., 2015; Hu et al., 2016), the ratio of PM₁ to TEOM-determined
133 PM_{2.5} in this work appeared to be a bit higher. The difference may be due to: (1) the contribution of semi-volatile species to
134 PM_{2.5} varied greatly among different periods and different locations, because TEOM is heated to 50 °C during the measurement,
135 which might have caused significant losses of semi-volatile species, e.g., ammonium nitrate and semi-volatile organics; and
136 (2) the contribution of particles in the size range of 1-2.5 μm to the total PM_{2.5} might also change among different pollution
137 episodes and different sites.

138 **2.3 Positive Matrix Factorization of organic aerosol matrix**

139 To determine potential sources of OA, the ACSM mass spectra were processed using the ME-2 algorithm implemented
140 with the toolkit SoFi (Source Finder) developed by Canonaco et al. (2013). The so-called *a* value approach allows the user to
141 introduce a priori information in forms of known factor profiles or time series to obtain a rather unique solution and thus reduce
142 the rotational ambiguity of the PMF2 algorithm. The spectra and error matrices of organics were prepared according to the
143 protocol summarized by Ulbrich et al. (2009) and Zhang et al. (2011). Given the interferences of the internal standard of
144 naphthalene at *m/z* 127-129 and the low signal-to-noise ratio of larger ions, we only considered ions up to *m/z* 120 in this study.
145 A reference HOA profile, which is an average of multiple ambient data sets taken from Ng et al. (2011b), was introduced to
146 constrain the model performance with *a* value varying from 0 to 1. Following the guidelines presented by Canonaco et al.
147 (2013) and Crippa et al. (2014), an optimal solution involving four factors with *a* value of 0.1 was accepted. Detailed analyses
148 of the factor time series, mass spectra, diurnal patterns, and correlations with external tracers can be found in the supplementary
149 information (Fig. S2-S6). Note that before using the ME-2 engine, we also attempted to perform PMF analysis with the PMF2
150 algorithm for 1 to 8 factors. The solutions were thoroughly evaluated following the recommendations outlined in Zhang et al.
151 (2011) and the results of three- and four-factor solutions at $f_{\text{peak}}=0$ are shown in Fig. S7-S8. The three-factor solution indicates
152 the identification of a coal combustion OA (CCOA), a biomass burning OA (BBOA) and an oxygenated OA (OOA). But the
153 CCOA factor seems to be mixed with the signals from hydrocarbon-like components related to traffic emissions, which is
154 especially evident given the two noticeable peaks in the diurnal profile of the CCOA-factor during morning and evening rush
155 hours. In the four-factor solution, the additional factor could not be physically explained and showed indications of factor
156 splitting. Solutions with 5 to 8 factors show further splitting and mixing of factors. Our inability to separate an individual HOA
157 factor using the PMF2 algorithm is probably due to the minor contribution of traffic emissions in Handan, consistent with the
158 fact that the PMF2 algorithm tends to have difficulty in accurately retrieving minor factors (Ulbrich et al., 2009).

159 **3 Results and discussions**

160 **3.1 Overview of aerosol characteristics**

161 Frequent and persistent haze episodes were observed during the campaign, especially from December 16 to December
162 25, 2015, when an extremely polluted and long-lasting haze event occurred. Based on TEOM measurements, only 4 days met

163 the US National Ambient Air Quality Standards (NAAQS, 35 $\mu\text{g}/\text{m}^3$ for the 24 h average of $\text{PM}_{2.5}$) and 13 days met the Chinese
164 NAAQS Grade II (75 $\mu\text{g}/\text{m}^3$ for the 24 h average of $\text{PM}_{2.5}$) for the whole study period of 65 days. In other words, the daily
165 average $\text{PM}_{2.5}$ concentrations exceeded the US NAAQS and the Chinese NAAQS on 94% and 80% of the days, respectively
166 (Fig. 2). On December 22, the daily $\text{PM}_{2.5}$ concentration reached the highest value of 725.7 $\mu\text{g}/\text{m}^3$, leading to the first “red”
167 haze alarm (http://www.cma.gov.cn/kppd/kppdsytj/201310/t20131028_229921.html) ever in Hebei Province. The
168 meteorological conditions were stagnant with calm winds throughout the study period (WS usually less than 1.5 m/s), although
169 relatively high WS (generally > 1.5 m/s) with cleaner air from northwest of Handan occasionally interrupted the haze evolution
170 process (Fig. 2b). The RH varied from 11.7% to 94.8%, generally with higher values for more polluted periods and lower
171 values during cleaner periods. No precipitation occurred throughout the entire campaign.

172 Hourly PM_{10} concentrations fluctuated dramatically from 4.2 $\mu\text{g}/\text{m}^3$ to 700.8 $\mu\text{g}/\text{m}^3$ (Fig. 2g). The average PM_{10}
173 concentration was 187.6 $\mu\text{g}/\text{m}^3$, more than twice as high as that observed in the well-known severe haze event that occurred in
174 Beijing in January 2013 (Sun et al., 2014; Zhang et al., 2014). Organics constituted a major fraction of PM_{10} , contributing 45%
175 on average during this study, followed by sulfate (15%), nitrate (14%), and ammonium (12%). The large fraction of organics
176 in PM_{10} was comparable to previous observations in other areas of NCP during wintertime (Sun et al., 2013a; Zhang et al.,
177 2013; Huang et al., 2014). The average chloride contribution (9%) is relatively high compared to that previously observed in
178 NCP region. Submicron nonrefractory chloride in the aerosol phase can be directly emitted from different sources (e.g.,
179 biomass burning and coal combustion) (Lobert et al., 1999; McCulloch et al., 1999), and formed in the atmosphere through
180 gas-to-particle conversion (e.g., NH_4Cl partitioning) (Baek et al., 2006). Considering that chloride demonstrated pronouncedly
181 enhanced peaks during night and it showed good correlations with CCOA and BBOA ($r=0.72$ and 0.80 , respectively), a large
182 fraction of chloride during wintertime was thought to be from primary emissions at night. On average, BC accounted for 5%
183 of total PM_{10} . Its distinct peaks at morning and evening rush hours suggested that BC was mainly associated with traffic
184 emissions. In the daytime, PM_{10} was dominated by secondary species because of active photochemistry, whereas the
185 contributions of primary species were significantly increased at night, probably caused by enhanced primary emissions from
186 fuel combustion coupled with shallow boundary layer height (Fig. S9).

187 Ambient CO is an indicator for the intensities of anthropogenic emissions. The hourly CO concentration was as high as
188 10 ppm during the study period, higher than those observed in other areas of China (Andreae et al., 2008; Quan et al., 2014;
189 Yang et al., 2015). Interestingly, the temporal pattern of organics tracked well with that of CO ($r = 0.84$, Fig.2), implying that
190 combustion emissions were a significant source of organic aerosols in Handan, i.e., traffic, coal combustion, and biomass
191 burning. In addition, during severe haze episodes with high NO_x and CO concentrations, ozone remained at nearly zero for
192 several days instead of showing a regular diurnal variation, indicating active ozone titration by NO and a strong influence of
193 primary emissions on haze pollution in this study.

194 3.2 Source apportionment of organic aerosol

195 In this study, three POA factors (HOA, BBOA and CCOA) and one SOA factor (OOA) were resolved by analyzing the

196 ACSM OA mass spectra using the ME-2 algorithm. OOA was the largest contributor to OA mass with an average fraction of
197 39% (Fig. 3). The traffic-related HOA only accounted for 7% of total OA, which was in accordance with the fact that PMF
198 analysis performed with the PMF2 algorithm had difficulty to retrieve it (see Sect. 2.3 for more details). On average, primary
199 sources dominated the OA mass (61%) during this winter study, consistent with the results from previous winter studies in the
200 NCP region (Sun et al., 2013a; Zhang et al., 2014; Hu et al., 2016; Sun et al., 2016a). The discussion below focuses on the
201 characteristics, sources and processes of each OA factor.

202 **3.2.1 Hydrocarbon-like OA**

203 The HOA factor shows a mass spectrum highly similar to those of freshly emitted traffic or other fossil combustion
204 aerosols (Zhang et al., 2005a; Lanz et al., 2007; Li et al., 2016a). Its profile is dominated by alkyl fragment signatures, the
205 $C_nH_{2n+1}^+$ (m/z 29, 43, 57) and $C_nH_{2n-1}^+$ (m/z 27, 41, 55) ion series. The time series of HOA correlated well with those of NO_x
206 and BC ($r = 0.75$ and 0.74 , respectively, Fig. 3e), two tracers of vehicle emissions. The diurnal pattern of HOA (Fig. 3i) further
207 confirmed the association of HOA with traffic activities, as it showed two obvious peaks during morning and evening rush
208 hours. On average, HOA only accounted for 7% of total OA in Handan; a much smaller fraction than observed in the nearby
209 megacities of Beijing and Tianjin (Sun et al., 2013a; Wang et al., 2015). The small HOA fraction in this study is consistent
210 with findings from a previous source apportionment study which revealed that transportation was a minor source of
211 atmospheric particles in Handan (Wang et al., 2014). Bivariate polar plots, which present the concentrations of air pollutants
212 as a function of WS and WD using the OpenAir software (Carslaw and Ropkins, 2012), demonstrated higher concentrations
213 of HOA under relatively low WS ($<1.5\text{m/s}$), suggesting that HOA was substantially influenced by local emission sources, in
214 accordance with its primary characteristics (Fig. S10).

215 **3.2.2 Coal combustion OA**

216 Although coal combustion has rarely been reported as an important source of organic aerosols in the US or Europe, it is
217 a large emitter of organics in China (Cao et al., 2006). According to Zhang et al. (2008b), organic carbon can contribute up to
218 70% of emitted $PM_{2.5}$ for different types of coal combustion in China. During wintertime, coal is the primary fuel for various
219 industries (e.g. power generation, steel milling, and cement production) as well as residential heating in the NCP region. Thus
220 a considerable contribution from coal combustion to OA concentration was expected in this study. Compared to HOA and
221 BBOA, the mass spectrum of CCOA showed strong signals at higher m/z , especially a significant peak at m/z 115, and the
222 temporal trend of CCOA correlated tightly with that of m/z 115 ($r=0.99$, Fig. 3). These findings are similar to observations
223 made in Beijing, Changdao, Xi'an, and Lanzhou during winter, where OA factors representing coal combustion were
224 determined (Hu et al., 2013; Elser et al., 2016; Sun et al., 2016a; Xu et al., 2016). Further, a recent study by Zhou et al. (2016)
225 has shown that the ACSM mass spectra of OA from residential coal combustion emissions tend to present a high peak at m/z
226 115. In addition, CCOA was also found to correlate relatively well with chloride ($r = 0.72$) during this study, consistent with
227 the fact that coal combustion is also an important emission source of chloride.

228 Figure 4 compares the OA composition in this study with those of previous winter studies in China. During wintertime,
229 CCOA was observed to contribute a significant fraction of the fine PM mass in regions to the north of the Yangtze River (e.g.
230 Beijing, Lanzhou, and Handan), due to domestic coal combustion for heating in winter. However, little to no CCOA was
231 observed in areas located to the south of the Yangtze River, e.g. Nanjing, Jiaxing, and Ziyang, mainly reflecting the lack of
232 central heating provided by the Chinese government in this region during winter. In this study, similar to the results observed
233 in Beijing and Lanzhou (Sun et al., 2013a; Hu et al., 2016; Xu et al., 2016), CCOA on average accounted for 29% of total OA,
234 with a minimum of 13% at noon and a maximum of 32% at midnight. However, the average mass concentration of CCOA in
235 Handan ($23.1 \mu\text{g}/\text{m}^3$) was much higher than those observed in previous studies. Given the high consumption of coal and the
236 important role of coal combustion for aerosol pollution in Handan, control of air pollutant emissions from coal combustion
237 through technology renewal is essential for air quality improvement in this area.

238 3.2.3 Biomass burning OA

239 Biomass burning, including wildfires, forest and agricultural burning, and domestic biofuel combustion, is one of the
240 largest emission sources of organics worldwide (Ramanathan et al., 2001). Biomass burning releases air pollutants that have
241 adverse effects on respiratory organs and reduce lung function of human beings (Regalado et al., 2006). In the NCP region,
242 during the harvest seasons in summer and autumn with open agricultural burning, biomass burning is a major influence on
243 aerosol loadings and characteristics. For example, at a suburban site near Beijing during summertime, Sun et al. (2016b)
244 observed that the contribution of BBOA to OA increased from 6% during the non-biomass burning period to 21% during the
245 biomass burning period. During wintertime, as most previous studies of this region were performed in the megacity of Beijing
246 where coal combustion dominates the energy consumption, BBOA was seldom resolved or found to be a minor fraction of
247 total OA mass (Sun et al., 2013a; Zhang et al., 2014; Huang et al., 2014; Sun et al., 2016a). However, for many small and
248 medium-sized cities in the NCP region, domestic combustion of wood and crop residuals for cooking and home heating is very
249 popular in the countryside during wintertime and could emit large amounts of air pollutants (Zhang et al., 2008a; Ding et al.,
250 2012). In Hebei province, biomass burning accounted for 52% of primary organic carbon emissions during the winter of 2015
251 according to the Multi-resolution Emission Inventory for China (MEIC; <http://meicmodel.org/>).

252 In this study, a BBOA factor with high mass concentrations was clearly observed, the mass spectrum of which was
253 characterized by the prominent peaks at m/z 60 and 73, two indicative tracers of biomass burning (Alfarra et al., 2007; Aiken
254 et al., 2009; Lee et al., 2010). The time series of the BBOA varied dramatically and correlated well with that of CO ($r = 0.72$),
255 which was mainly emitted from combustion-related sources. BBOA showed clear diurnal variations with low mass
256 concentrations occurring during daytime and high mass concentrations arising at night. Consistent with the emission inventory,
257 BBOA on average accounted for 25% of total OA mass, with an average concentration of $20.7 \mu\text{g}/\text{m}^3$, much higher than that
258 observed in other areas of China during wintertime (Fig. 4), indicating the important role of biomass burning emissions in
259 aerosol pollution in Handan. Polar plots showed that high BBOA concentrations were mainly related to local emissions (Fig.
260 S10), probably associated with cooking and residential heating using biofuel.

261 3.2.4 Oxygenated OA

262 Although two or more OOA factors with different oxidation degree and formation pathways have been resolved in
263 previous wintertime studies in China (Xu et al., 2015; Sun et al., 2016a), only one OOA factor was observed in our study. The
264 mass spectrum of OOA presented a pattern similar to those reported before (e.g., Zhang et al., 2005b; Ng et al., 2011c) with a
265 prominent peak at m/z 44 (15.8% of the total OOA signal). In addition, OOA showed a temporal trend similar to those of
266 sulfate and nitrate, and correlated strongly with the sum of secondary inorganic species (SIA = sulfate + nitrate + ammonium)
267 (Fig. 5). The polar plots of OOA and secondary inorganic species exhibited similar spatial distributions, with high concentration
268 hotspots located in the northeast, especially during polluted periods (Fig. S10). The temporal variation profile of OOA was
269 much different from those of the POA factors ($r = 0.50$; Fig. 5). As shown in Fig. 3, while POA varied dramatically between
270 day and night due to the influence of local emissions, the mass concentrations of OOA often built up gradually and remained
271 at high levels for several days until being swept away by clean air masses. These results are consistent with OOA being
272 representative of SOA. Although the diurnal profile of OOA was overall flat in this study, the mass fraction of OOA to total
273 OA increased significantly during daytime, reaching a maximum of 64% at 14:00 BJT (Fig. 3n).

274 3.3 Diurnal variations and insights into aerosol sources

275 3.3.1 Weekdays versus weekends

276 As air pollutants are mainly emitted from anthropogenic sources in Handan, comparing the diurnal profiles of aerosol
277 species between weekdays and weekends would provide insights into the variations of different emission sources and
278 atmospheric processes. Generally speaking, weekdays span Monday to Friday, whereas weekends include Saturday and Sunday.
279 However, because the physical and chemical processes in the atmosphere are not completed instantaneously, the variations of
280 aerosol species may be influenced by the carry-over effect of the previous day. Thus, we alternatively define weekdays from
281 Tuesday to Friday, and weekends only including Sunday. With this classification, differences in the diurnal variations between
282 weekdays and weekends are more visible. Comparisons of the diurnal cycles using the Monday-Friday and Saturday-Sunday
283 definitions are presented in the supplementary information (Fig. S11).

284 As displayed in Fig. 6, the diurnal variations of meteorological parameters did not significantly change from weekdays
285 to weekends, providing a good opportunity to investigate the influence of anthropogenic activities. As expected, the diurnal
286 pattern of HOA, which is associated with traffic emissions, presented a more distinct morning peak on weekdays. This was
287 also the case for BC, CO, and NO_x, which are all fossil fuel combustion tracers. However, the evening rush hour peaks of these
288 species did not show much of a difference between weekdays and weekends, indicating that human activities in the evening
289 were not significantly reduced on weekends. Other aerosol species showed generally similar diurnal trends for weekdays and
290 weekends, similar to the results observed in Beijing (Sun et al., 2013a). In contrast, stronger weekday vs. weekend differences
291 were observed in the US, where the mass concentrations of aerosol species are obviously lower during weekends (Young et
292 al., 2016; Zhou et al., 2016). Results from the present study reveal that active anthropogenic emissions tend to persist

293 throughout the entire week in polluted regions in Handan, leading to limited differences in the concentrations and compositions
294 of major air pollutants between weekdays and weekends. The exception is traffic emissions, for which the morning rush hour
295 peak is more prominent during weekdays.

296 **3.3.2 Polluted versus non-polluted periods**

297 To gain further insights into the evolution of aerosol particles throughout the day, especially during hazy conditions, we
298 explored the diurnal differences of meteorological conditions and air pollutants between polluted and non-polluted days (Fig.
299 7). According to the CNAAQs Grade II of daily $PM_{2.5}$ concentrations ($75 \mu\text{g}/\text{m}^3$), only the 13 days (out of a total of 65 days)
300 were found to meet the requirement and are considered to be non-polluted in this study; the rest are defined as polluted periods.
301 Note that of these 13 non-polluted days, only 3 days achieved the 24 h CNAAQs Grade I level of $PM_{2.5}$ ($35 \mu\text{g}/\text{m}^3$).

302 The temperature was relatively low throughout the period, averaging 2.1°C and 0.2°C on polluted and non-polluted days,
303 respectively. The RH during polluted periods was slightly higher during daytime, favoring the aqueous-phase processing of
304 atmospheric pollutants. The influence of RH is discussed in detail in Sect. 3.5. Stagnant weather conditions with lower wind
305 speeds were observed on polluted days, especially during nighttime, which would aggravate the accumulation of aerosol
306 pollution. Unsurprisingly, the mass concentrations of aerosol components and the mixing ratios of gaseous species were much
307 higher on polluted days. But the diurnal differences between polluted and non-polluted periods could provide some information
308 regarding their evolutionary processes. The diurnal profiles of secondary inorganic species (i.e. sulfate, nitrate, and
309 ammonium), were flatter on polluted days. For example, in the diurnal profile of nitrate during polluted periods, the maximum
310 and minimum concentrations were different by only 13 % or $4.4 \mu\text{g}/\text{m}^3$. This behavior is consistent with the comparison of
311 polar plots between polluted and non-polluted days, which indicated a significant effect of regional transport on polluted
312 periods for secondary species. In contrast, the diurnal trends of primary aerosol species, e.g. HOA, BBOA, and CCOA, during
313 polluted periods differed substantially from those during non-polluted periods. Compared to the relatively flat diurnal profiles
314 on non-polluted days, the mass concentrations of HOA, BBOA, and CCOA were strongly enhanced at nighttime on polluted
315 days. This suggests that the sharp increases of primary species at night, especially those of BBOA and CCOA, may play an
316 important role in haze formation.

317 **3.4 Evolution of aerosol characteristics with increasing PM_{10} concentration**

318 Identifying the responsible emission sources and formation pathways during haze events is essential to effectively
319 implement emission controls, especially with the increased frequency of haze events during winter. In this study, the whole
320 period is divided into polluted and non-polluted days, as described in Sect. 3.3.2. The average PM_{10} concentration during
321 polluted days ($211 \mu\text{g}/\text{m}^3$) was more than three times higher than that during non-polluted days ($49 \mu\text{g}/\text{m}^3$). However, the
322 average aerosol composition did not show obvious changes between these two types of days, indicating the synchronous
323 increase of all aerosol species (Fig. 8a). Indeed, during polluted days, the average mass concentrations of all aerosol species,
324 except for BC, were approximately four times as high as those during non-polluted days (Fig. 8b). Sulfate, CCOA, and BBOA

325 showed the highest polluted/non-polluted ratios, which were 5.3, 5.0, and 5.5, respectively (Fig. 8b). Given the higher average
326 RH on polluted days (average $\pm 1\sigma = 56.5\pm 18.8\%$) compared to non-polluted days (average $\pm 1\sigma = 40.9\pm 18.7\%$), aqueous-phase
327 processing likely has increased the production of sulfate (Wang et al., 2012; Zheng et al., 2015; Elser et al., 2016). During
328 polluted days, the average oxidation ratio of sulfur (molar ratio of sulfate to sum of sulfate and SO₂) was 0.27, higher than that
329 on non-polluted days (0.16). On the other hand, the strong increases of CCOA and BBOA were possibly caused by enhanced
330 gas-to-particle partitioning associated with high PM mass loadings during polluted periods (Mader et al., 2002). Interestingly,
331 compared to aerosol species, CO showed a lower polluted/non-polluted ratio of approximately 2. A possible reason is that CO
332 has a longer atmospheric lifetime compared to aerosol particles, thus it has a more elevated regional background concentration.
333 Note that the polluted/non-polluted ratios for SO₂ and NO_x were also lower compared to the aerosol species. This is potentially
334 a result of enhanced aqueous phase oxidation of SO₂ and NO_x as well as more efficient wet deposition, since the more polluted
335 periods were generally more humid.

336 Figure 9 further displays the average hourly variations of the mass fractions of aerosol species as a function of PM₁
337 concentration. The nitrate fraction went up a bit and then showed a decreasing trend with increasing PM₁ mass loading whereas
338 the contribution of sulfate increased from 12% to 20% as PM₁ concentration developed from 100 $\mu\text{g}/\text{m}^3$ to 600 $\mu\text{g}/\text{m}^3$. Since
339 it was unlikely that the emission sources of the main gaseous precursors of these two species (i.e., NO_x and SO₂) had changed
340 significantly during our study, the observed changes in aerosol compositions suggest different formation mechanisms of nitrate
341 and sulfate during wintertime. The substantially elevated production of sulfate during high PM episodes was likely attributable
342 to higher ambient RH, which facilitated sulfate production through aqueous-phase reactions of SO₂ (Kim et al., 2016; Li et al.,
343 2016b; Sun et al., 2013b). The oxidation ratio of sulfur increased from 0.1 to 0.4 when PM₁ concentration raised from ~ 10
344 $\mu\text{g}/\text{m}^3$ to 600 $\mu\text{g}/\text{m}^3$. The mass fractions of different OA factors varied widely as PM₁ concentrations increased. The
345 contribution of HOA to total PM₁ was minor and remained relatively stable across all mass loadings. However, the mass
346 fractions of CCOA and BBOA increased nearly linearly with PM₁ concentrations rising from ~ 20 $\mu\text{g}/\text{m}^3$ to 300 $\mu\text{g}/\text{m}^3$ and
347 plateaued at higher aerosol loadings. OOA, a surrogate of SOA, showed the opposite PM-loading dependency, and its
348 contribution decreased slightly with increasing PM₁ concentration. The study of Sun et al. (2013a) in Beijing also found a
349 growing contribution of CCOA and a declining contribution of OOA with increasing PM₁ concentrations during wintertime.
350 These results reveal the important role of POA in the development of high PM pollution during wintertime. Indeed, the scatter
351 plot of OA vs. PM₁ concentrations (Fig. 9c) demonstrates that higher mass fractions of organics in PM₁ were associated with
352 elevated POA contributions to total OA, especially when PM₁ concentrations were more than 200 $\mu\text{g}/\text{m}^3$ (Fig. 9c). Overall, the
353 results here suggest that secondary formation of sulfate (mainly from SO₂ emitted by coal combustion), and primary emissions
354 of organics from coal combustion and biomass burning are important factors driving the development of winter haze pollution
355 in Handan.

356 3.5 A case study on an intense haze episode and the influence of meteorological conditions

357 From December 14 to December 28, 2015, an extremely severe haze episode occurred and was characterized with a steady

358 build-up of air pollutants, including fine particles and CO, over a period of ~ 5 days (Dec. 17 -21, 2015) followed by
359 approximately 4 days of heavy air pollution, during which the average CO mixing ratio was 6.7 ppm and the average PM₁
360 concentration was 500.1 µg/m³ (Fig. 10). This episode ended on Dec. 25, during which winds from the northwest brought in
361 cleaner air, leading to dramatic reductions of air pollutants. This type of evolutionary process has been frequently observed in
362 Beijing during autumn and winter, and is called “sawtooth cycles” by Jia et al. (2008). In this study, the whole haze cycle was
363 divided into five stages: (1) a clean period (Stage 1), (2) an almost linear increasing period of PM₁ concentration (Stage 2), (3)
364 a remarkably high pollution period lasting for four days (Stage 3), (4) an abruptly cleaned up period (Stage 4), and (5) another
365 clean period as the start of a new cycle (Stage 5). As shown in Fig. 10, each stage was initiated by a sudden change in the WD
366 and air masses from different regions via the HYSPLIT back trajectories (Draxler and Rolph, 2013). This indicates that
367 meteorological changes are important driving forces during the evolution of haze episodes.

368 Stage 1 was characterized by high winds from the northwest, which brought clean air masses from Western Siberia.
369 Aerosols associated with this air mass origin were largely free of high anthropogenic emissions and appeared to be aged with
370 a high contribution of secondary species. Consistently, the CO concentration during stage 1 was relatively low. During stage
371 2, the WD changed and the WS was lower than 1 m s⁻¹. The air masses from the northern and southern areas of Handan were
372 influenced by high anthropogenic emissions in northern Hebei and Henan province, respectively. Thus, the PM₁ concentration
373 steadily increased during this stage, with an average of 164.6 µg/m³. Stage 3 was dominated by southerly and northerly winds
374 and really stagnant conditions with low WS. On December 23, air masses from the southern and northern areas of Hebei
375 circulated around Handan, leading to the accumulation of air pollutants including PM and CO. The average PM₁ concentration
376 during stage 3 was 500.1 µg/m³, with the hourly maximum reaching as high as 700.8 µg/m³, much higher than that observed
377 during the severe haze episode in Beijing in January 2013 (~300 µg/m³; Sun et al., 2014). Accompanied with a high CO
378 concentration (average of ~7 ppm) during stage 3, O₃ concentration remained at a very low level of almost zero and with
379 minimal diurnal variations, suggesting that gas-phase oxidation might not be a dominant mechanism for haze formation.
380 Moreover, stage 3 was characterized with high RH, exceeding 70% most of the time, which would promote the aqueous-phase
381 formation of secondary species. Indeed, a high mass fraction of secondary species, especially a notable increase in sulfate
382 contribution, was observed during stage 3. During stage 4, due to the return of cleaner air masses long transported from
383 northwest, air pollutant concentrations in Handan decreased dramatically and PM₁ concentration decreased from 443.7 µg/m³
384 to 34.1 µg/m³ within only 12 hours.

385 To further evaluate the influence of air mass origins on aerosol characteristics, we performed the cluster analysis of
386 HYSPLIT back trajectories for the whole study period to elucidate the relationship between aerosol concentration or
387 composition and different clusters. As shown in Fig. S12, the whole NCP region was heavily polluted, with high PM₁
388 concentrations for all four clusters. Overall, the aerosol compositions were similar among different clusters. However, we
389 indeed observed an important role played by winds in altering aerosol characteristics according to the above case study.
390 Referring to the haze cycle analysis, we attempted to apply another classification method based on WD and WS. Periods with
391 WS exceeding 1.5 m s⁻¹ from the northwest of Handan were denoted as “NW_HWS”, whereas the remaining periods were

392 classified as “Others” (Fig. 11). As expected, the PM₁ concentration of “Others” was more than six times higher than that of
393 “NW_HWS”. Secondary aerosol species (i.e. sulfate, nitrate, ammonium and OOA) contributed 66% of total PM₁ for
394 “NW_HWS”. As air masses associated with “Others” were more strongly influenced by anthropogenic sources, the main
395 primary species (i.e. HOA, BBOA, CCOA and BC), accounted for a higher fraction of 32% for “Others”. These results
396 highlight the importance of high winds from the northwest of Handan in alleviating PM levels and changing aerosol
397 composition during wintertime.

398 As mentioned previously, the sulfate contribution during stage 3 was visibly enhanced under high RH, revealing the
399 effects of RH on aerosol processing. Many previous studies have observed the increased production of secondary inorganic
400 aerosol species through aqueous-phase processing. In this study, we used the oxidation ratios of sulfur and nitrogen, defined
401 as $f_S = nSO_4^{2-} / (nSO_4^{2-} + nSO_2)$ and $f_N = nNO_3^- / (nNO_3^- + nNO_x)$, respectively, to explore the influence of RH on
402 aerosol formation (Fig. 12). Under relatively dry conditions (RH<50%), both f_S and f_N were almost constant. However,
403 when RH>50%, f_S started to increase linearly, similar to the results observed by Zheng et al. (2015) in Beijing. In comparison,
404 f_N showed a small increase at RH 60-70% and then decreased a bit when RH>90%, suggesting different roles of aqueous-
405 phase reactions in the formation of sulfate and nitrate. Recently, studies of aqueous-phase chemistry have paid increasing
406 attention to organic components. Ge et al. (2012) observed the strong enhancement of SOA during a fog event in the Central
407 Valley of California during winter. Based on high resolution mass spectra from an AMS, Sun et al. (2016a) retrieved an
408 aqueous-phase-processed SOA (aq-OOA) that tracked well with RH in Beijing during wintertime. However, the mass fraction
409 of SOA in total OA in this study remained relatively stable and showed no dependency on RH (Fig. 12c). The RH-binned bulk
410 composition of submicron aerosol also only exhibited an obvious increase of sulfate at high RH (Fig. S13). One explanation
411 for this observation is that the variations of SOA contribution may be largely interfered by high fractions of POA across
412 different RH values. Another explanation is that, a portion of OOA formed through aqueous phase reactions may be
413 incorporated into fog droplets, which are too large to be transmitted into the ACSM aerodynamic lens, as reported by Ge et al.
414 (2012). This explanation is consistent with the results obtained by studying a fog event in London, in which no increase in
415 OOA concentration was detected by AMS measurement, whereas the single particle mass spectrometry observed aqueous-
416 phase SOA production (Dall’Osto et al., 2009).

417 **4 Conclusions**

418 To characterize aerosol sources and formation processes under high anthropogenic emissions in the NCP region, a field
419 campaign was conducted in Handan during the extremely polluted winter of 2015-2016. For the entire study period, only 13
420 out of 65 days met the Chinese NAAQS Grade II of 75 µg/m³ for daily PM_{2.5}. The average concentration of submicron aerosol
421 was 187.6 µg/m³, with hourly values fluctuating dramatically by a factor of ~ 150, from 4.2 µg/m³ to 700.8 µg/m³. Organics
422 dominated the bulk composition of submicron aerosols (44.6% of PM₁ mass), similar to previous observations in the NCP
423 region during wintertime. PMF analysis identified three primary sources of organic aerosol, i.e. traffic, coal combustion and

424 biomass burning, and one SOA factor. CCOA was the largest contributor to POA, on average accounting for 29%, followed
425 by BBOA (25%). The mass fraction of HOA in total OA was only 7%, indicating the minor contribution of traffic emissions
426 in Handan. Although the aerosol concentration during polluted days was more than three times higher than that during non-
427 polluted days, little variation was observed in the average aerosol bulk composition, revealing the relatively synchronous
428 increase of all aerosol species during haze evolution. Stagnant weather conditions, with low wind speed and high RH, and
429 strong enhancement of primary species at nighttime, prompted haze formation during polluted days. Variation of aerosol mass
430 fractions with hourly increasing PM₁ concentration further revealed that secondary formation of sulfate (mainly from SO₂
431 emitted by coal combustion) and primary emissions from coal combustion and biomass burning, are important factors driving
432 haze formation. This is mainly related to large emissions of air pollutants from coal and biomass combustion during wintertime,
433 especially for simple household stoves with low combustion efficiency. Overall, sulfate, chloride, and CCOA on average
434 accounted for a total of 37% of PM₁ mass (Fig. 1c), showing the important role of coal combustion in air pollution in Handan.
435 Given the continuing high consumption of coal for various industries and residential heating in winter, technology-based
436 emission controls on coal combustion would effectively improve the air quality in Handan.

437 A severe haze episode that started with a steady build-up of aerosol pollution followed by an abrupt clean period was
438 studied. Our results indicate the strong influence of meteorological conditions on haze evolution. With high anthropogenic
439 emissions around Handan, the whole study region was heavily polluted. However, high aerosol loadings can be rapidly
440 alleviated by strong winds from the northwestern. Under high RH (RH>50%), the oxidation ratio of sulfur increased linearly,
441 suggesting the important role of aqueous phase chemistry in sulfate formation during wintertime. Results from this study
442 provide useful insights into aerosol chemistry and haze evolution in Hebei province during wintertime, and have important
443 implications for pollution control in this heavily polluted area.

444 **Acknowledgements**

445 This work was funded by the National Natural Science Foundation of China (41571130035 and 41625020) and the Ford
446 Company. Haiyan Li was partially supported by the Doctoral Short-Term Visiting-Abroad Foundation of Tsinghua University,
447 Beijing. Qi Zhang acknowledges the Changjiang Scholars program of the Chinese Ministry of Education. We also give special
448 acknowledgement to lab members in the Department of Environmental Engineering, Hebei University of Engineering, Handan,
449 China, whose help was invaluable in setting up this field campaign.

450 **References**

451 Aiken, A. C., Salcedo, D., Cubison, M. J., Huffman, J. A., DeCarlo, P. F., Ulbrich, I. M., Docherty, K. S., Sueper, D., Kimmel,
452 J. R., Worsnop, D. R., Trimborn, A., Northway, M., Stone, E. A., Schauer, J. J., Volkamer, R. M., Fortner, E., de Foy, B.,
453 Wang, J., Laskin, A., Shutthanandan, V., Zheng, J., Zhang, R., Gaffney, J., Marley, N. A., Paredes-Miranda, G., Arnott, W.

454 P., Molina, L. T., Sosa, G., and Jimenez, J. L.: Mexico City aerosol analysis during MILAGRO using high resolution aerosol
455 mass spectrometry at the urban supersite (T0) – Part 1: Fine particle composition and organic source apportionment, *Atmos.*
456 *Chem. Phys.*, 9, 6633–6653, doi:10.5194/acp-9-6633-2009, 2009.

457 Alfara, M. R., Prévôt, A. S. H., Szidat, S., Sandradewi, J., Weimer, S., Lanz, V. A., Schreiber, D., Mohr, M., and Baltensperger,
458 U.: Identification of the mass spectral signature of organic aerosols from wood burning emissions, *Environ. Sci. Technol.*, 41,
459 5770– 5777, 2007.

460 Andreae, M. O., Schmid, O., Yang, H., Chand, D., Yu, J. Z., Zeng, L. M., and Zhang, Y. H.: Optical properties and chemical
461 composition of the atmospheric aerosol in urban Guangzhou, China, *Atmos. Environ.*, 42, 6335–6350, 2008.

462 Baek, B. H., Koziel, J., and Aneja, V. P.: A preliminary review of gas-to-particle conversion, monitoring, and modeling efforts
463 in the USA. *Int. J. Global Environ. Iss.*, 6 (2/3), 204–230, 2006. Canonaco, F., Crippa, M., Slowik, J. G., Baltensperger, U., and
464 Prévôt, A. S. H.: SoFi, an Igor based interface for the efficient use of the generalized multilinear engine (ME-2) for source
465 apportionment: application to aerosol mass spectrometer data, *Atmos. Meas. Tech.*, 6 3649–3661, doi:10.5194/amt-6-3649-
466 2013, 2013.

467 Cao, G., Zhang, X., and Zheng, F.: Inventory of black carbon and organic carbon emissions from China, *Atmos. Environ.*, 40,
468 6516–6527, 2006.

469 Carslaw, D. C. and Ropkins, K.: openair - An R package for air quality data analysis, *Environmental Modelling & Software*,
470 27-500 28, 52-61, 2012.

471 Crippa, M., Canonaco, F., Lanz, V. A., Äijälä, M., Allan, J. D., Carbone, S., Capes, G., Ceburnis, D., Dall'Osto, M., Day, D.
472 A., DeCarlo, P. F., Ehn, M., Eriksson, A., Freney, E., Hildebrandt Ruiz, L., Hillamo, R., Jimenez, J. L., Junninen, H., Kiendler-
473 Scharr, A., Kortelainen, A.-M., Kulmala, M., Laaksonen, A., Mensah, A. A., Mohr, C., Nemitz, E., O'Dowd, C., Ovadnevaite,
474 J., Pandis, S. N., Petäjä, T., Poulain, L., Saarikoski, S., Sellegri, K., Swietlicki, E., Tiitta, P., Worsnop, D. R., Baltensperger, U.,
475 and Prévôt, A. S. H.: Organic aerosol components derived from 25 AMS data sets across Europe using a consistent ME-2 based
476 source apportionment approach, *Atmos. Chem. Phys.*, 14, 6159– 6176, doi:10.5194/acp-14-6159-2014, 2014.

477 Dall'Osto, M., Harrison, R. M., Coe, H., and Williams, P.: Real-time secondary aerosol formation during a fog event in London,
478 *Atmos. Chem. Phys.*, 9, 2459-2469, doi:10.5194/acp-9-2459-2009, 2009.

479 Ding, J., Zhong, J., Yang, Y., Li, B., Shen, G., Su, Y., Wang, C., Li, W., Shen, H., Wang, B., Wang, R., Huang, Y., Zhang, Y.,
480 Cao, H., Zhu, Y., Simonich, S.L., and Tao, S.: Occurrence and exposure to polycyclic aromatic hydrocarbons and their
481 derivatives in a rural Chinese home through biomass fuelled cooking, *Environ. Pollut.*, 169, 160–166, 2012.

482 Draxler, R. R. and Rolph, G.D.: HYSPLIT (Hybrid single-particle lagrangian integrated trajectory) Model access via NOAA
483 ARL READY Website, available at: <http://www.arl.noaa.gov/HYSPLIT.php> (last access: 2014), NOAA Air Resources
484 Laboratory, College Park, MD, 2013.

485 Elser, M., Huang, R.-J., Wolf, R., Slowik, J. G., Wang, Q., Canonaco, F., Li, G., Bozzetti, C., Daellenbach, K. R., Huang, Y.,
486 Zhang, R., Li, Z., Cao, J., Baltensperger, U., El-Haddad, I., and Prévôt, A. S. H.: New insights into PM_{2.5} chemical
487 composition and sources in two major cities in China during extreme haze events using aerosol mass spectrometry, *Atmos.*

488 Chem. Phys., 16, 3207-3225, doi:10.5194/acp-16-3207-2016, 2016.

489 Ge, X. L., Zhang, Q., Sun, Y. L., Ruehl, C. R., and Setyan, A.: Effect of aqueous-phase processing on aerosol chemistry and
490 size distributions in Fresno, California, during wintertime, Environ. Chem., 9, 221–235, doi:10.1071/EN11168, 2012.

491 Guo, S., Hu, M., Zamora, M. L., Peng, J. F., Shang, D. J., Zheng, J., Du, Z. F., Wu, Z., Shao, M., Zeng, L. M., Molina, M. J.,
492 and Zhang, R. Y.: Elucidating severe urban haze formation in China, P Natl Acad Sci USA, 111, 17373-17378,
493 10.1073/pnas.1419604111, 2014.

494 Hu, W. W., Hu, M., Yuan, B., Jimenez, J. L., Tang, Q., Peng, J. F., Hu, W., Shao, M., Wang, M., Zeng, L. M., Wu, Y. S., Gong,
495 Z. H., Huang, X. F., and He, L. Y.: Insights on organic aerosol aging and the influence of coal combustion at a regional receptor
496 site of central eastern China, Atmos. Chem. Phys., 13, 10095-10112, doi:10.5194/acp-13-10095-2013, 2013.

497 Hu, W., Hu, M., Hu, W., Jimenez, J. L., Yuan, B., Chen, W., Wang, M., Wu, Y., Chen, C., Wang, Z., Peng, J., Zeng, L., and
498 Shao, M.: Chemical composition, sources and aging process of submicron aerosols in Beijing: contrast between summer and
499 winter, J. Geophys. Res., 121, 1955–1977, doi:10.1002/2015JD024020, 2016.

500 Huang, R.-J., Zhang, Y., Bozzetti, C., Ho, K.-F., Cao, J.-J., Han, Y., Daellenbach, K. R., Slowik, J.G., Platt, S. M., Canonaco,
501 F., Zotter, P., Wolf, R., Pieber, S. M., Bruns, E. A., Crippa, M., Ciarelli, G., Piazzalunga, A., Schwikowski, M., Abbaszade, G.,
502 SchnelleKreis, J., Zimmermann, R., An, Z., Szidat, S., Baltensperger, U., El Haddad, I., and Prévôt, A. S. H.: High secondary
503 aerosol contribution to particulate pollution during haze events in China, Nature, 514, 218–222, doi:10.1038/nature13774,
504 2014.

505 Jia, Y. T., Rahn, K. A., He, K. B., Wen, T. X., and Wang, Y. S.: A novel technique for quantifying the regional component of
506 urban aerosol solely from its sawtooth cycles, J. Geophys. Res., 113, D21309, doi:10.1029/2008JD010389, 2008. Kim, H.,
507 Zhang, Q., Bae, G.-N., Kim, J. Y., and Lee, S. B.: Sources and atmospheric processing of wintertime aerosols in Seoul, Korea:
508 Insights from real-time measurements using a high-resolution aerosol mass spectrometer, Atmos. Chem. Phys. Discuss.,
509 doi:10.5194/acp-2016-855, in review, 2016.

510 Lanz, V. A., Alfarra, M. R., Baltensperger, U., Buchmann, B., Hueglin, C., and Prévôt, A. S. H.: Source apportionment of
511 submicron organic aerosols at an urban site by factor analytical modelling of aerosol mass spectra, Atmos. Chem. Phys., 7,
512 1503-1522, doi:10.5194/acp-7-1503-2007, 2007.

513 Lee, T., Sullivan, A. P., Mack, L., Jimenez, J. L., Kreidenweis, S. M., Onasch, T. B., Worsnop, D. R., Malm, W., Wold, C. E.,
514 Hao, W. M., and Collett, J. L.: Variation of chemical smoke marker emissions during flaming vs. smoldering phases of
515 laboratory open burning of wildland fuels, Aerosol Sci. Technol., 44, 1–5, doi:10.1080/02786826.2010.499884, 2010.

516 Li, H., Zhang, Q., Duan, F., Zheng, B., and He, K.: The “Parade Blue”: effects of short-term emission control on aerosol
517 chemistry, Faraday Discuss., 189, 317-335, 2016a.

518 Li, H., Duan, F., He, K., Ma, Y., Kimoto, T., Huang, T.: Size-dependent characterization of atmospheric particles during winter
519 in Beijing, Atmosphere, 7, 2016b.

520 Lobert, J. M., Keene, W. C., Logan, J. A., and Yevich, R.: Global chlorine emissions from biomass burning: Reactive Chlorine
521 Emissions Inventory, J Geophys Res-Atmos, 104, 8373-8389, Doi 10.1029/1998jd100077, 1999.

522 Mader, B. T., and Pankow, J. F.: Study of the effects of particle-phase carbon on the gas/particle partitioning of semivolatile
523 organic compounds in the atmosphere using controlled field experiments. *Environ. Sci. Technol.*, 36, 5218-5228, 2002.

524 McCulloch, A., Aucott, M. L., Benkovitz, C. M., Graedel, T. E., Kleiman, G., Midgley, P. M., and Li, Y. F.: Global emissions
525 of hydrogen chloride and chloromethane from coal combustion, incineration and industrial activities: Reactive Chlorine
526 Emissions Inventory, *J Geophys Res-Atmos*, 104, 8391-8403, Doi 10.1029/1999jd900025, 1999.

527 Middlebrook, A. M., Bahreini, R., Jimenez, J. L., and Canagaratna, M. R.: Evaluation of Composition-Dependent Collection
528 Efficiencies for the Aerodyne Aerosol Mass Spectrometer using Field Data, *Aerosol Sci. Tech.*, 46, 258–271, 2012.

529 Ng, N. L., Canagaratna, M. R., Zhang, Q., Jimenez, J. L., Tian, J., Ulbrich, I. M., Kroll, J. H., Docherty, K. S., Chhabra, P. S.,
530 Bahreini, R., Murphy, S. M., Seinfeld, J. H., Hildebrandt, L., Donahue, N. M., DeCarlo, P. F., Lanz, V. A., Prévôt, A. S. H.,
531 Dinar, E., Rudich, Y., and Worsnop, D. R.: Organic aerosol components observed in Northern Hemispheric datasets from
532 Aerosol Mass Spectrometry, *Atmos. Chem. Phys.*, 10, 4625-4641, doi:10.5194/acp-10-4625-2010, 2010.

533 Ng, N. L., Herndon, S. C., Trimborn, A., Canagaratna, M. R., Croteau, P., Onasch, T. M., Sueper, D., Worsnop, D. R., Zhang,
534 Q., Sun, Y. L., and Jayne, J. T.: An Aerosol Chemical Speciation Monitor (ACSM) for routine monitoring of atmospheric
535 aerosol composition, *Aerosol Sci. Technol.*, 45, 770–784, 2011a.

536 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Zhang, Q., Ulbrich, I. M., and Worsnop, D. R.: Real-time methods for estimating
537 organic component mass concentrations from aerosol mass spectrometer data, *Environ. Sci. Technol.*, 45, 910–916, 2011b.

538 Ng, N. L., Canagaratna, M. R., Jimenez, J. L., Chhabra, P. S., Seinfeld, J. H., and Worsnop, D. R.: Changes in organic aerosol
539 composition with aging inferred from aerosol mass spectra, *Atmos. Chem. Phys.*, 11, 6465-6474, doi:10.5194/acp-11-6465-
540 2011, 2011c.

541 Paatero, P. and Tapper, U.: Positive Matrix Factorization: a non-negative factor model with optimal utilization of error
542 estimates of data values, *Environmetrics*, 5, 111–126, 1994.

543 Paatero, P.: The multilinear engine - A table-driven, least squares program for solving multilinear problems, including the n-
544 way parallel factor analysis model, *J. Comput. Graph. Stat.*, 8, 854– 888, 1999.

545 Petzold, A. and Schonlinner, M.: Multi-angle absorption photometry – a new method for the measurement of aerosol light
546 absorption and atmospheric black carbon, *J. Aerosol Sci.*, 35, 421–441, 2004.

547 Petzold, A., Schloesser, H., Sheridan, P. J., Arnott, W. P., Ogren, J. A., and Virkkula, A.: Evaluation of multiangle absorption
548 photometry for measuring aerosol light absorption, *Aerosol Sci. Tech.*, 39, 40–51, 2005.

549 Pope III, C. A. and Dockery, D. W.: Health Effects of Fine Particulate Air Pollution: Lines that Connect, *J. Air Waste Manage.*
550 56, 709–742, 2006.

551 Pöschl, U.: Atmospheric Aerosols: Composition, Transformation, Climate and Health Effects, *Angew. Chem. Int. Ed.*, 44,
552 7520–7540, 2005.

553 Quan, J., Tie, X., Zhang, Q., Liu, Q., Li, X., Gao, Y., and Zhao, D.: Characteristics of heavy aerosol pollution during the 2012–
554 2013 winter in Beijing, China, *Atmos. Environ.*, 88, 83–89, doi:10.1016/j.atmosenv.2014.01.058, 2014.

555 Ramanathan, V., Crutzen, P. J., Kiehl, J. T., and Rosenfeld, D.: Aerosols, climate and the hydrological cycle, *Science*, 294,
556 2119–2124, 2001.

557 Regalado J, Pérez-Padilla R, Sansores R, Ramirez JIP, Brauer M, Paré P, and Vedal, S.: The effect of biomass burning on
558 respiratory symptoms and lung function in rural Mexican women. *Am J Respir Crit Care Med.*, 174, 901–905, 2006.

559 Seinfeld, J. H. and Pandis, S. N.: *Atmospheric Chemistry and Physics: From Air Pollution to Climate Change*, John Wiley &
560 Sons, New York, 2nd edition, 1232 pp., ISBN-13: 978-0-471-72018-8, 2006.

561 Sun, J. Y., Zhang, Q., Canagaratna, M. R., et al.: Highly time- and size-resolved characterization of submicron aerosol particles
562 in Beijing using an Aerodyne Aerosol Mass Spectrometer, *Atmos. Environ.*, 44, 131–140, 2010.

563 Sun, Y., Wang, Z., Dong, H., Yang, T., Li, J., Pan, X., Chen, P., and Jayne, J. T.: Characterization of summer organic and
564 inorganic aerosols in Beijing, China with an Aerosol Chemical Speciation Monitor, *Atmos. Environ.*, 51, 250–259,
565 doi:10.1016/j.atmosenv.2012.01.013, 2012.

566 Sun, Y. L., Wang, Z. F., Fu, P. Q., Yang, T., Jiang, Q., Dong, H. B., Li, J., and Jia, J. J.: Aerosol composition, sources and
567 processes during wintertime in Beijing, China, *Atmos. Chem. Phys.*, 13, 4577–4592, doi:10.5194/acp-13-4577-2013, 2013a.

568 Sun, Y. L., Wang, Z. F., Fu, P. Q., Jiang, Q., Yang, T., Li, J., and Ge, X. L.: The impact of relative humidity on aerosol
569 composition and evolution processes during wintertime in Beijing, China, *Atmos. Environ.*, 77, 927–934,
570 doi:10.1016/j.atmosenv.2013.06.019, 2013b.

571 Sun, Y., Jiang, Q., Wang, Z., Fu, P., Li, J., Yang, T., and Yin, Y.: Investigation of the Sources and Evolution Processes of Severe
572 Haze Pollution in Beijing in January 2013, *J. Geophys. Res.*, 119, 4380–4398, 2014.

573 Sun, Y. L., Wang, Z. F., Du, W., Zhang, Q., Wang, Q. Q., Fu, P. Q., Pan, X. L., Li, J., Jayne, J., and Worsnop, D. R.: Long-term
574 real-time measurements of aerosol particle composition in Beijing, China: seasonal variations, meteorological effects, and
575 source analysis, *Atmos. Chem. Phys.*, 15, 10149–10165, doi:10.5194/acp-15-10149-2015, 2015.

576 Sun, Y., Du, W., Fu, P., Wang, Q., Li, J., Ge, X., Zhang, Q., Zhu, C., Ren, L., Xu, W., Zhao, J., Han, T., Worsnop, D. R., and
577 Wang, Z.: Primary and secondary aerosols in Beijing in winter: sources, variations and processes, *Atmos. Chem. Phys.*, 16,
578 8309–8329, doi:10.5194/acp-16-8309-2016, 2016a.

579 Sun, Y., Jiang, Q., Xu Y., Ma, Y., Zhang, Y., Liu, X., Li, W., Wang, F., Li, J., Wang, P., and Li, Z.: Aerosol characterization
580 over the North China Plain: Haze life cycle and biomass burning impacts in summer, *J. Geophys. Res. Atmos.*, 121, 2508–
581 2521, doi:10.1002/2015JD024261, 2016b.

582 Takegawa, N., Miyakawa, T., Kuwata, M., et al.: Variability of submicron aerosol observed at a rural site in Beijing in the
583 summer of 2006, *J. Geophys. Res.*, 114, D00G05, doi:10.1029/2008JD010857, 2009.

584 Ulbrich, I. M., Canagaratna, M. R., Zhang, Q., Worsnop, D. R., and Jimenez, J. L.: Interpretation of organic components from
585 Positive Matrix Factorization of aerosol mass spectrometric data, *Atmos. Chem. Phys.*, 9, 2891–2918, doi:10.5194/acp-9-2891-
586 2009, 2009.

587 Wang, G., Cheng, S., Li, J., Lang, J., Wen, W., Yang, X., and Tian, L.: Source apportionment and seasonal variation of PM_{2.5}
588 carbonaceous aerosol in the Beijing–Tianjin–Hebei Region of China. *Environ. Monit. Assess.* 187, 2015.

589 Wang, L. T., Wei, Z., Yang, J., Zhang, Y., Zhang, F. F., Su, J., Meng, C. C., and Zhang, Q.: The 2013 severe haze over southern
590 Hebei, China: model evaluation, source apportionment, and policy implications, *Atmos. Chem. Phys.*, 14, 3151-3173,
591 doi:10.5194/acp-14-3151-2014, 2014.

592 Wang, X., Wang, W., Yang, L., Gao, X., Nie, W., Yu, Y., Xu, P., Zhou, Y., and Wang, Z.: The secondary formation of
593 inorganic aerosols in the droplet mode through heterogeneous aqueous reactions under haze conditions, *Atmos. Environ.*, 63,
594 68–76, 2012.

595 Wei, Z., Wang, L.T., Chen, M.Z., and Zheng, Y.: The 2013 severe haze over the Southern Hebei, China: PM_{2.5} composition
596 and source apportionment, *Atmos. Pollut. Res.* 5, 759-768, 2014.

597 Xu, J., Shi, J., Zhang, Q., Ge, X., Canonaco, F., Prévôt, A. S. H., Vonwiller, M., Szidat, S., Ge, J., Ma, J., An, Y., Kang, S.,
598 and Qin, D.: Wintertime organic and inorganic aerosols in Lanzhou, China: Sources, processes and comparison with the results
599 during summer, *Atmos. Chem. Phys. Discuss.*, doi:10.5194/acp-2016-278, in review, 2016.

600 Xu, L., Suresh, S., Guo, H., Weber, R. J., and Ng, N. L.: Aerosol characterization over the southeastern United States using
601 high-resolution aerosol mass spectrometry: spatial and seasonal variation of aerosol composition and sources with a focus on
602 organic nitrates, *Atmos. Chem. Phys.*, 15, 7307-7336, doi:10.5194/acp-15-7307-2015, 2015.

603 Yang, Y. R., Liu, X. G., Qu, Y., An, J. L., Jiang, R., Zhang, Y. H., Sun, Y. L., Wu, Z. J., Zhang, F., Xu, W. Q., and Ma, Q. X.:
604 Characteristics and formation mechanism of continuous hazes in China: a case study during the autumn of 2014 in the North
605 China Plain, *Atmos. Chem. Phys.*, 15, 8165-8178, doi:10.5194/acp-15-8165-2015, 2015.

606 Young, D. E., Kim, H., Parworth, C., Zhou, S., Zhang, X., Cappa, C. D., Seco, R., Kim, S., and Zhang, Q.: Influences of
607 emission sources and meteorology on aerosol chemistry in a polluted urban environment: results from DISCOVER-AQ
608 California, *Atmos. Chem. Phys.*, 16, 5427-5451, doi:10.5194/acp-16-5427-2016, 2016.

609 Zhang, J. K., Sun, Y., Liu, Z. R., Ji, D. S., Hu, B., Liu, Q., and Wang, Y. S.: Characterization of submicron aerosols during a
610 month of serious pollution in Beijing, 2013, *Atmos. Chem. Phys.*, 14, 2887-2903, doi:10.5194/acp-14-2887-2014, 2014.

611 Zhang, Q., Worsnop, D. R., Canagaratna, M. R., and Jimenez, J. L.: Hydrocarbon-like and oxygenated organic aerosols in
612 Pittsburgh: insights into sources and processes of organic aerosols, *Atmos. Chem. Phys.*, 5, 3289-3311, doi:10.5194/acp-5-
613 3289-2005, 2005a.

614 Zhang, Q., Alfarra, M. R., Worsnop, D. R., Allan, J. D., Coe, H., Canagaratna, M. R., and Jimenez, J. L.: Deconvolution and
615 quantification of hydrocarbon-like and oxygenated organic aerosols based on aerosol mass spectrometry, *Environ. Sci.*
616 *Technol.*, 39, 4938–4952, 2005b.

617 Zhang, Q., Jimenez, J. L., Canagaratna, M. R., Ulbrich, I. M., Ng, N. L., Worsnop, D. R., and Sun, Y.: Understanding
618 atmospheric organic aerosols via factor analysis of aerosol mass spectrometry: a review, *Anal. Bioanal. Chem.*, 401, 3045–
619 3067, 2011.

620 Zhang, R., Jing, J., Tao, J., Hsu, S.-C., Wang, G., Cao, J., Lee, C. S. L., Zhu, L., Chen, Z., Zhao, Y., and Shen, Z.: Chemical
621 characterization and source apportionment of PM_{2.5} in Beijing: seasonal perspective, *Atmos. Chem. Phys.*, 13, 7053-7074,
622 doi:10.5194/acp-13-7053-2013, 2013.

623 Zhang, Y., Dou, H., Chang, B., Wei, Z., Qiu, W., Liu, S., Liu, W., and Tao, S.: Emission of Polycyclic Aromatic Hydrocarbons
624 from Indoor Straw Burning and Emission Inventory Updating in China, *Annals of the New York Academy of Sciences*, 1140,
625 218–227, doi:10.1196/annals.1454.006, 2008a.

626 Zhang, Y., Schauer, J. J., Zhang, Y., Zeng, L., Wei, Y., Liu, Y., and Shao, M.: Characteristics of particulate carbon emissions
627 from real-world Chinese coal combustion, *Environ. Sci. Technol.*, 42, 5068–5073, 2008b.

628 Zhao, P. S., Dong, F., He, D., Zhao, X. J., Zhang, X. L., Zhang, W. Z., Yao, Q., and Liu, H. Y.: Characteristics of concentrations
629 and chemical compositions for PM_{2.5} in the region of Beijing, Tianjin, and Hebei, China, *Atmos. Chem. Phys.*, 13, 4631–4644,
630 doi:10.5194/acp-13-4631-2013, 2013.

631 Zheng, B., Zhang, Q., Zhang, Y., He, K. B., Wang, K., Zheng, G. J., Duan, F. K., Ma, Y. L., and Kimoto, T.: Heterogeneous
632 chemistry: a mechanism missing in current models to explain secondary inorganic aerosol formation during the January 2013
633 haze episode in North China, *Atmos. Chem. Phys.*, 15, 2031–2049, doi:10.5194/acp-15-2031-2015, 2015.

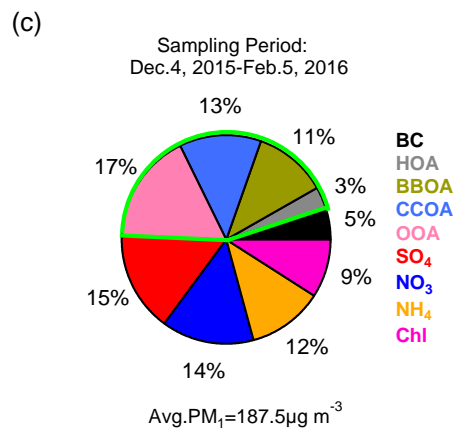
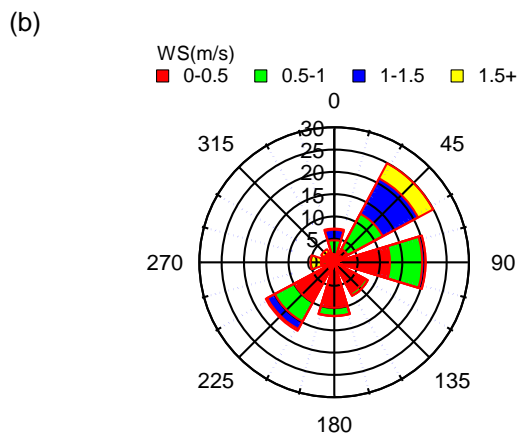
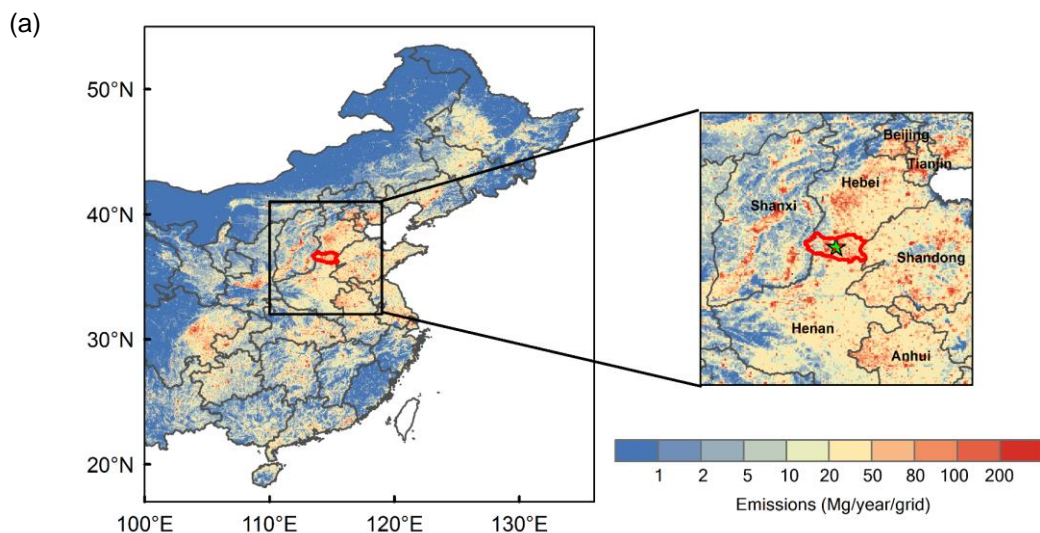
634 Zheng, G. J., Duan, F. K., Su, H., Ma, Y. L., Cheng, Y., Zheng, B., Zhang, Q., Huang, T., Kimoto, T., Chang, D., Pöschl, U.,
635 Cheng, Y. F., and He, K. B.: Exploring the severe winter haze in Beijing: the impact of synoptic weather, regional transport
636 and heterogeneous reactions, *Atmos. Chem. Phys.*, 15, 2969–2983, doi:10.5194/acp-15-2969-2015, 2015.

637 Zheng, G., Duan, F., Ma, Y., Zhang, Q., Huang, T., Kimoto, T., Cheng, Y., Su, H., and He, K.: Episode-Based Evolution Pattern
638 Analysis of Haze Pollution: Method Development and Results from Beijing, China, *Environ. Sci. Technol.*, 50, 4632–4641,
639 2016.

640 Zhou, S., Collier, S., Xu, J., Mei, F., Wang, J., Lee, Y. N., Sedlacek III, A. J., Springston, S. R., Sun, Y., and Zhang, Q.:
641 Influences of upwind emission sources and atmospheric processing on aerosol chemistry and properties at a rural location in
642 the Northeastern US, *J. Geophys. Res.*, 121, 6049–6065, doi:10.1002/2015JD024568., 2016.

643 Zhou, W., Jiang, J., Duan, L., and Hao, J.: Evolution of submicron organic aerosols during a complete residential coal
644 combustion process, *Environ. Sci. Technol.*, 50, 7861–7869, 2016.

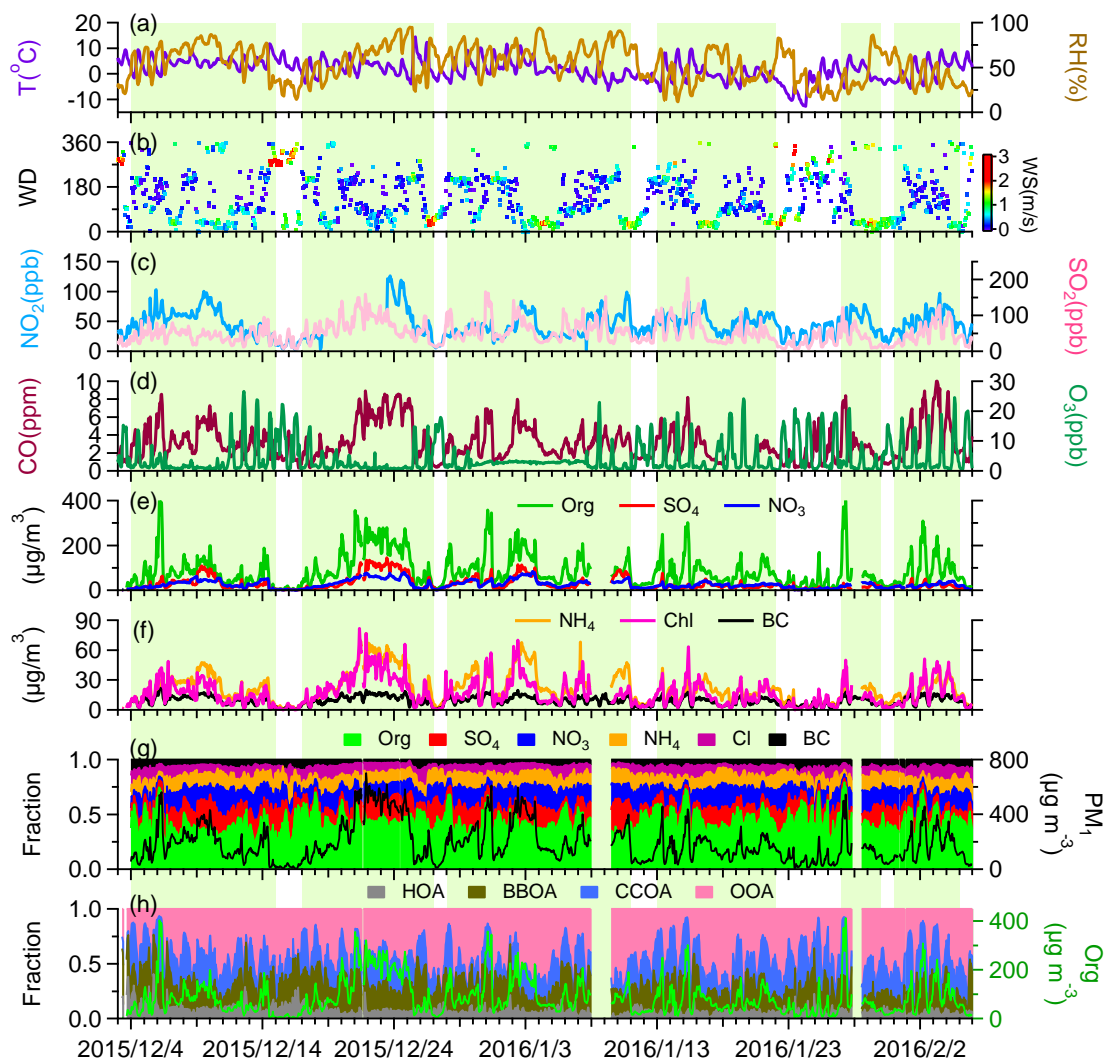
645



646

647 **Figure 1.** (a) Location of the sampling site in Handan in North China Plain. The map is color-coded by annual organic carbon emission
 648 rates modeled by Multi-resolution Emission Inventory for China (MEIC, <http://www.meicmodel.org>). The grid size is 0.05° × 0.05°. (b)
 649 Wind rose plot colored by wind speed for the entire period. Radial scales correspond to the frequency. (c) Compositional pie chart of
 650 submicron aerosol for the whole study, where the total organic fraction is outlined in green.

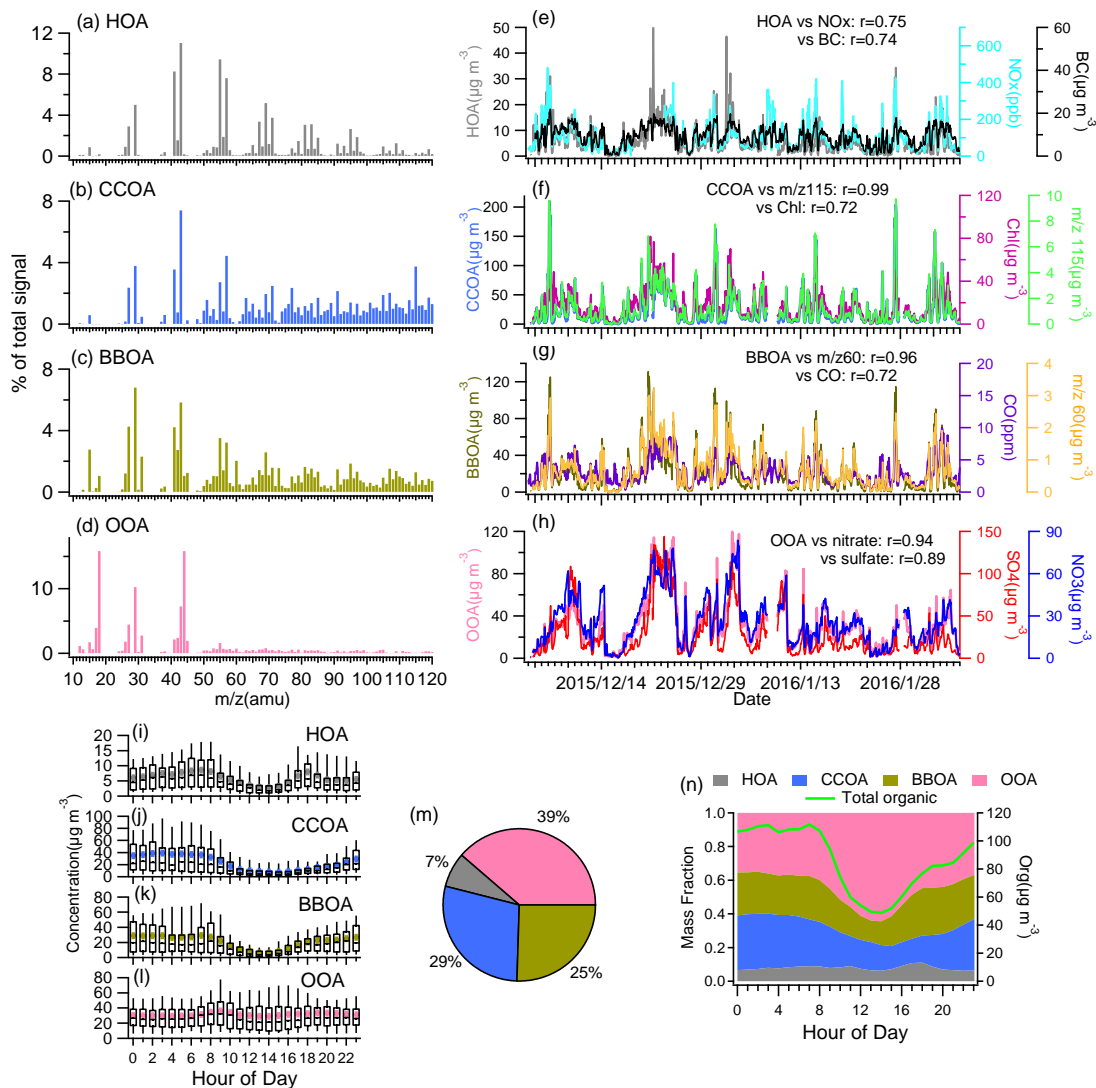
651



652

653 **Figure 2.** Time series of (a) ambient air temperature (T) and relative humidity (RH); (b) wind direction (WD) colored by wind speed
 654 (WS); (c) mixing ratios of NO₂ and SO₂; (d) mixing ratios of CO and O₃; (e) mass concentrations of organics, sulfate and nitrate; (f) mass
 655 concentrations of ammonium, chloride, and black carbon; (g) mass fractional contribution of chemical species to total PM₁ with the time
 656 series of total PM₁ concentration plotted in black on the right y-axis; (h) mass fractional contribution to total OA mass of the four factors
 657 derived from PMF analysis with the time series of organic aerosol plotted in green on the right y-axis. Days violating the CNAAQs for
 658 PM_{2.5} (= 75 µg m⁻³) are highlighted in pale green.

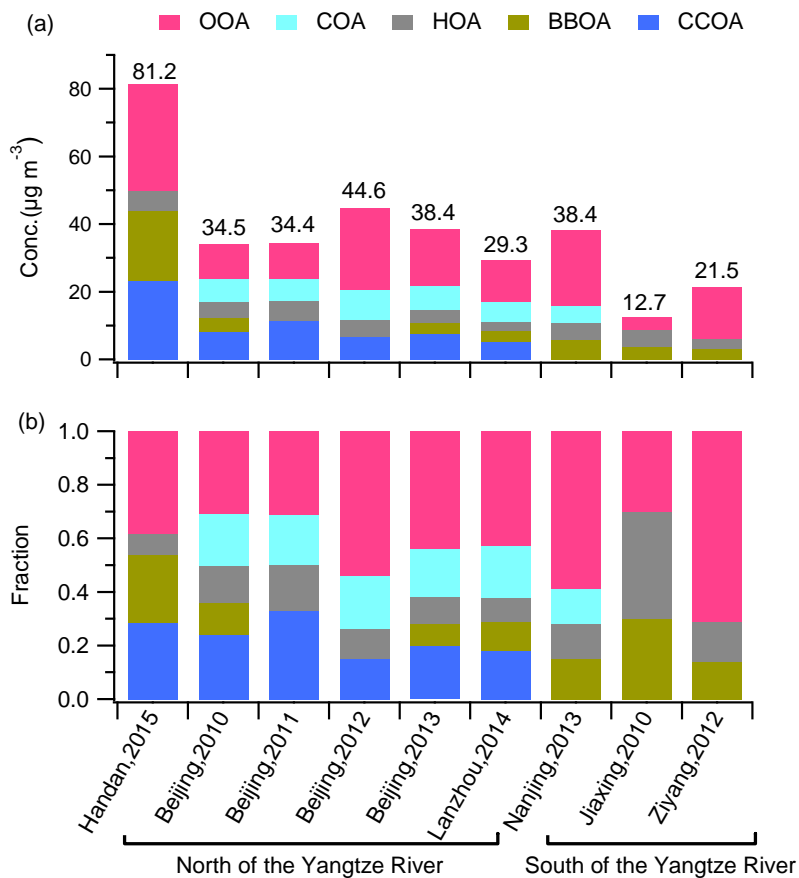
659



660

661 **Figure 3.** (a-d) Mass spectra of hydrocarbon-like OA (HOA), coal combustion OA (CCOA), biomass burning OA (BBOA), and
 662 oxygenated OA (OOA). (e-h) Time series of OA factors and the corresponding tracer compounds. (i-l) Diurnal patterns of OA factors. (m)
 663 Average fractional pie chart of OA factors to total OA for the campaign. (n) Average diurnal mass contributions of OA factors to total OA,
 664 with the average diurnal concentration of organics on the right y-axis.

665



666

667 **Figure 4.** Summary of the average (a) mass concentration and (b) chemical composition of organic aerosols from winter
 668 studies in China. The total concentration of OA ($\mu\text{g}/\text{m}^3$) is shown on the top of the bar in panel (a). See Table S1 in the
 669 Supplement for detailed information.

670

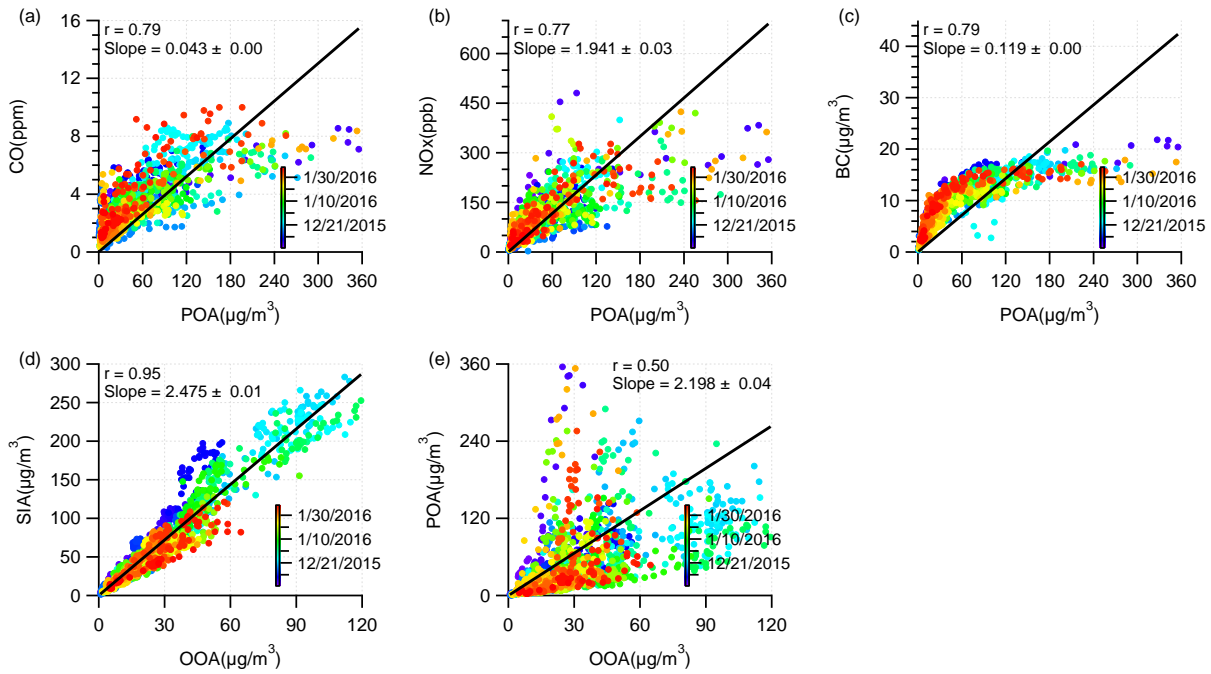
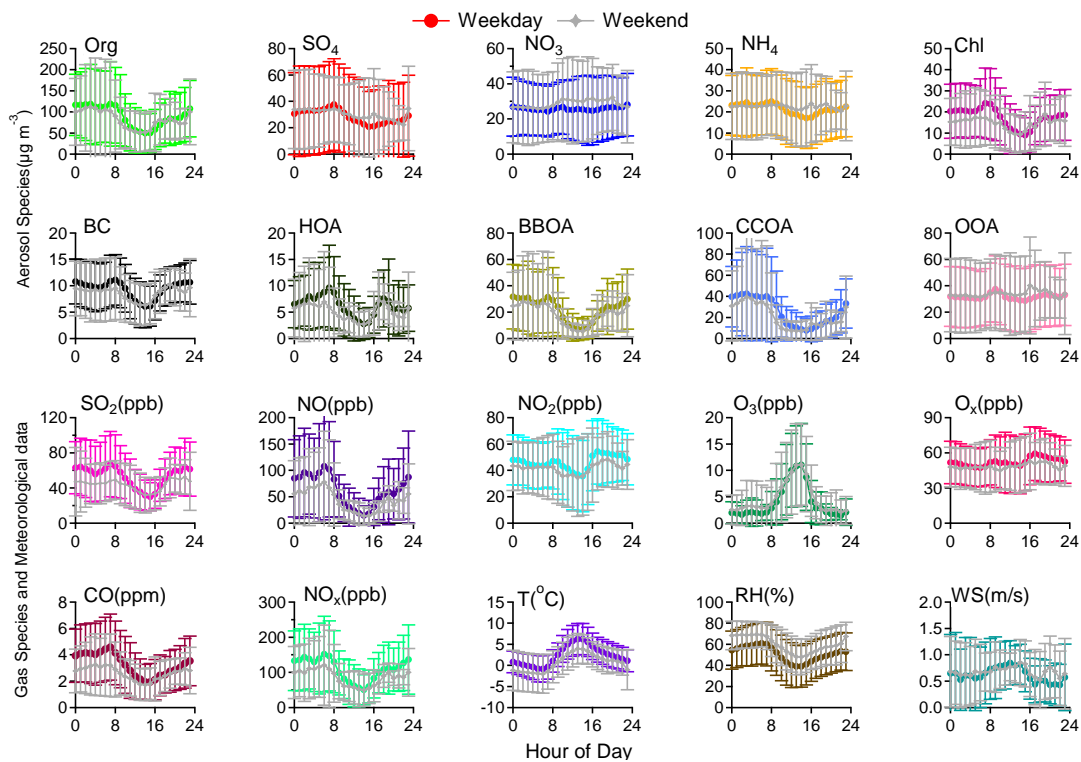


Figure 5. Scatter plots of (a) CO vs POA, (b) NO_x vs POA, (c) BC vs POA, (d) SIA vs OOA, and (e) POA vs OOA.

671
672

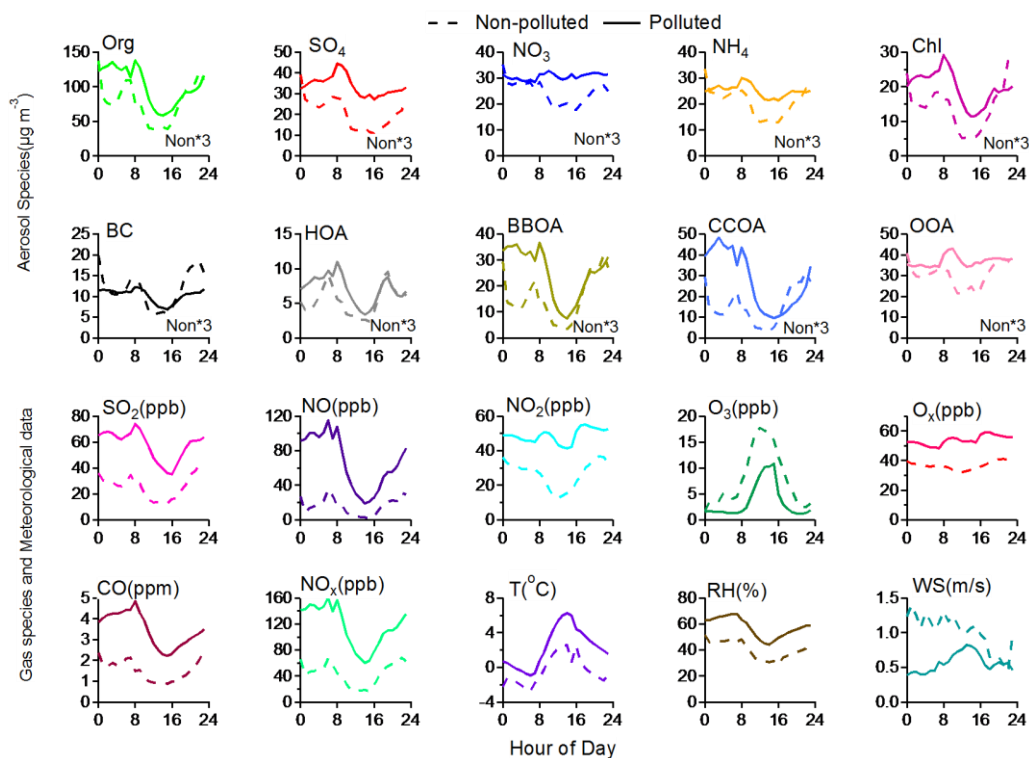
673



674

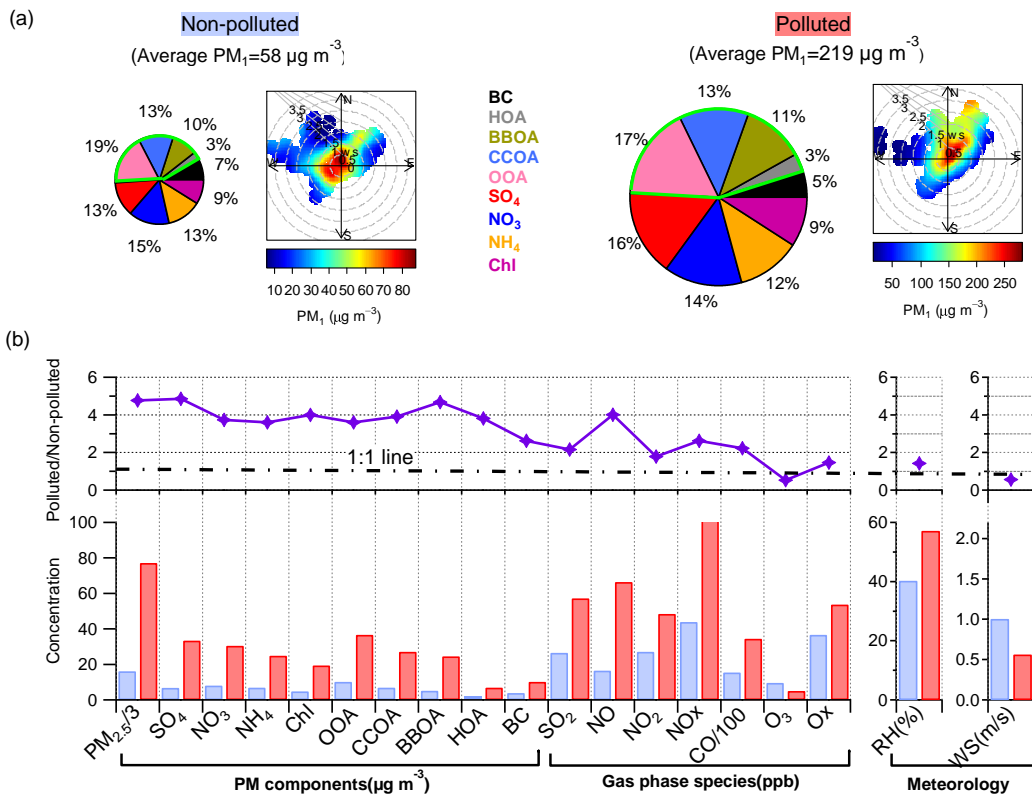
675 **Figure 6.** Average diurnal profiles along with the standard deviation of PM₁ species, four OA factors identified by PMF analysis, various
 676 gas-phase species, and meteorological parameters on weekdays (Tuesday to Friday inclusive) and weekends (Sunday only) during the
 677 campaign.

678



679
680
681
682

Figure 7. Average diurnal cycles of PM₁ species, four OA factors identified via PMF analysis, various gas-phase species, and meteorological parameters on polluted and non-polluted days. The mass concentrations of aerosol species during non-polluted periods are scaled by three factors in to highlight the differences in their diurnal trends on polluted and non-polluted days.



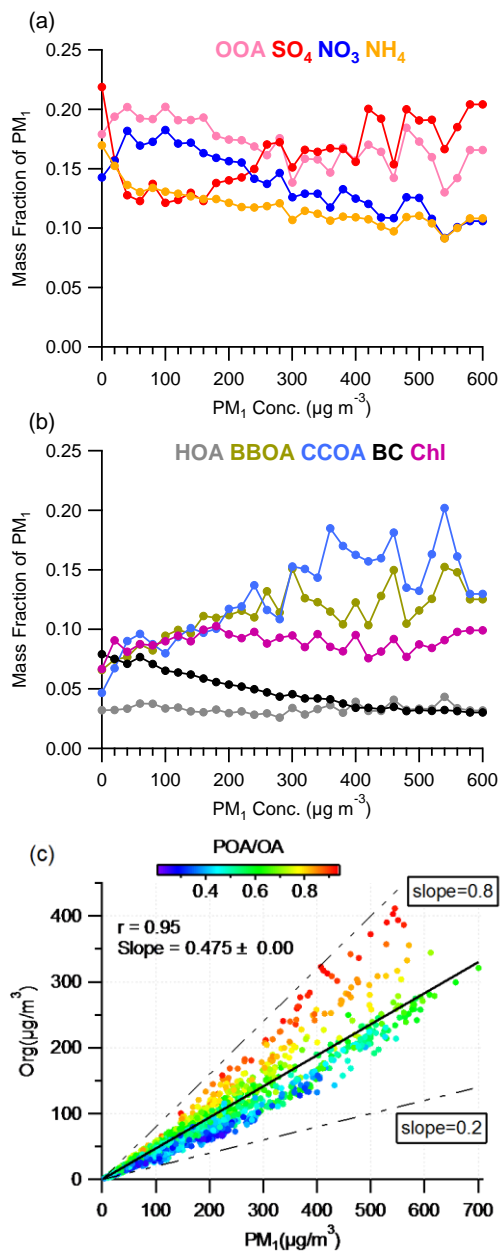
683

684

685

686

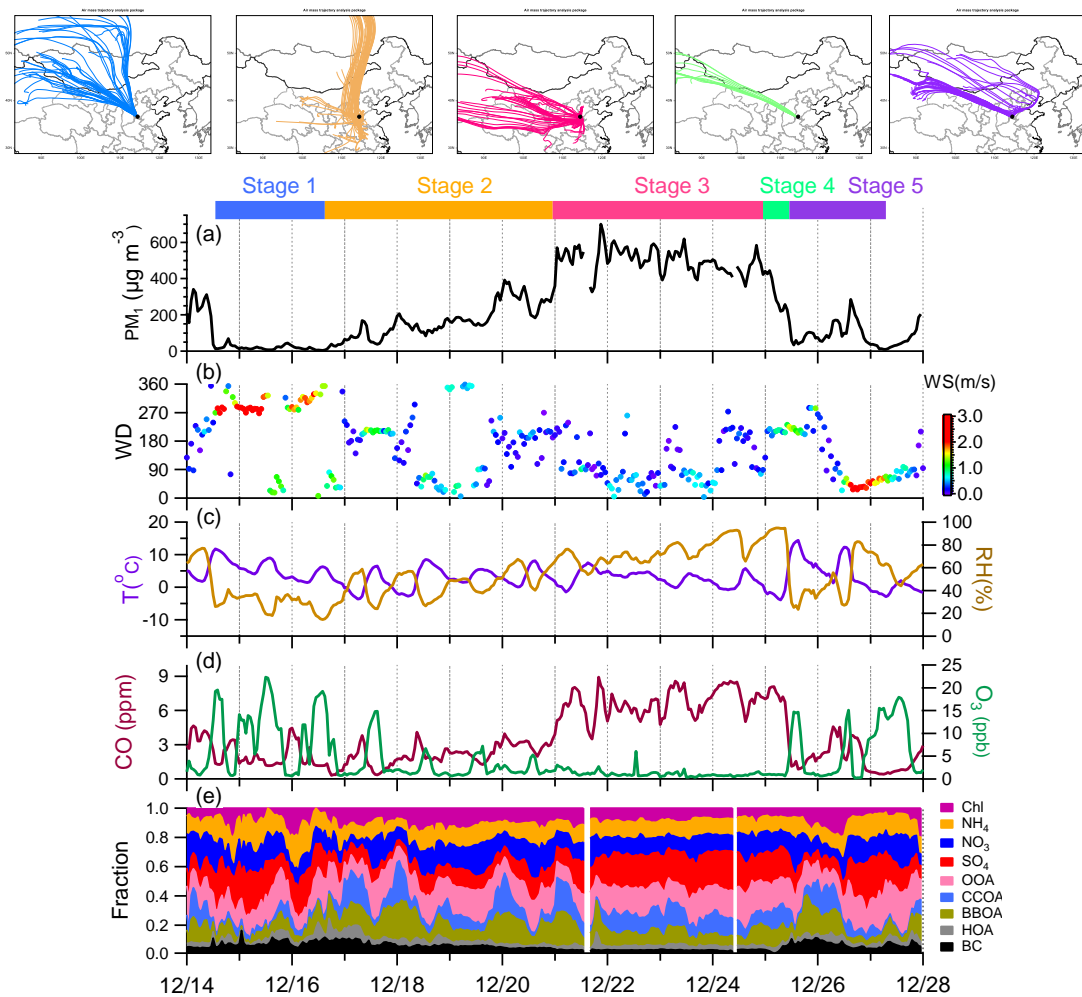
Figure 8. (a) Average PM_{10} composition and bivariate polar plots of PM_{10} concentration as a function of wind speed and wind direction for polluted and non-polluted periods. (b) Average concentration of PM_{10} components, gas phase species and average meteorological conditions during polluted and non-polluted days, with their polluted/non-polluted ratios shown in the top panel.



687

688 **Figure 9.** (a, b) Variations of the mass fractions of aerosol species as a function of PM₁ concentration. (c) Correlation plot of organics
 689 and PM₁ concentrations, colored by the mass fraction of POA in total OA.

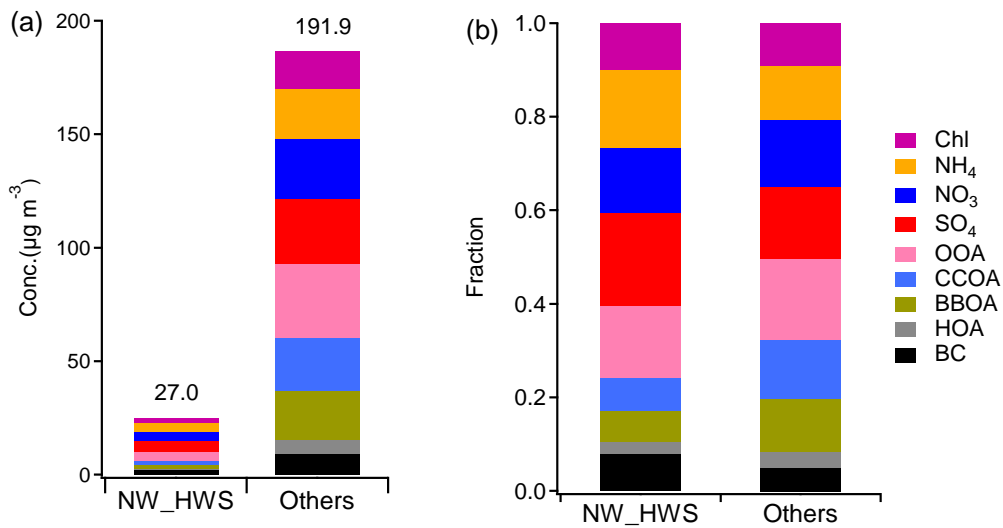
690



691

692 **Figure 10.** Evolution of (a) PM₁ concentration; (b) wind direction (WD) and wind speed (WS); (c) Temperature (T) and relative
 693 humidity (RH); (d) mixing ratios of CO and O₃; (e) mass fractions of aerosol species during a severe haze cycle from December 14 to
 694 December 28, 2016. The event was divided into five stages, with back trajectories of each stage shown on the top.

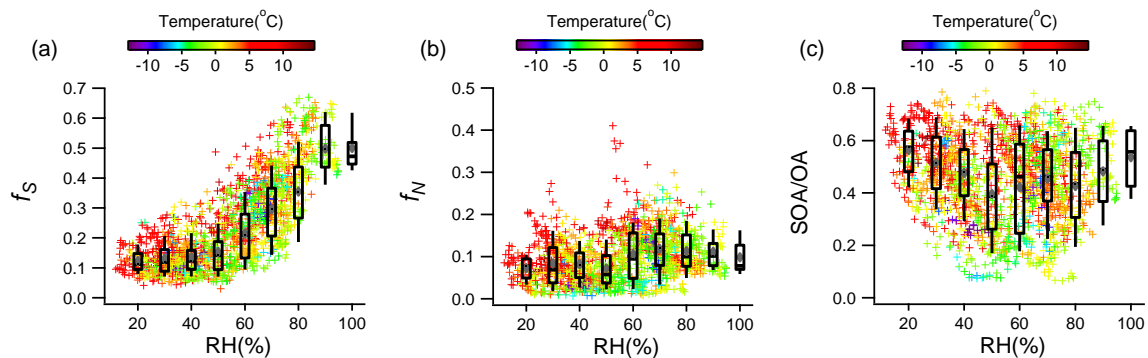
695



696

697 **Figure 11.** Comparisons of (a) mass concentrations of all PM₁ species and (b) fractional contributions of PM₁ species between
 698 "NW_HWS" and "Others" for the entire study period. The "NW_HWS" refers to high winds from the northwestern areas, and "Others"
 699 refers to the remaining.

700



701

702 **Figure 12.** Variations of (a) f_s , (b) f_N , and (c) the mass fraction of SOA in total OA plotted against increasing RH. The data are also
 703 binned according to RH values, and the mean (cross), median (horizontal line), 25th and 75th percentiles (lower and upper box), and 10th and
 704 90th percentiles (lower and upper whiskers) are shown for each bin.