Assessing the uncertainty of soil moisture impacts on convective precipitation using a new ensemble approach

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Abstract. Soil moisture amount and distribution controls evapotranspiration and thus impacts the occurrence of convective precipitation. Many recent model studies have proved this interaction and demonstrated that changes in initial soil moisture condition result in modified convective precipitation. At the same time, the chaotic behavior of the atmospheric system needs to be considered: any slight changes in the simulation setup, such as the chosen model domain, also result in modifications in the simulated precipitation field. This uncertainty caused by the stochastic variability is large compared to soil moisture effects as we will show. By shifting the model domain, an estimate on the uncertainty of the model results can be calculated. This uncertainty estimate, which includes ten simulations with shifted model boundaries, is compared to the effects on precipitation, which are caused by variations in soil moisture amount and local distribution. With this approach, the influence of soil moisture amount and distribution on convective precipitation is quantified. Deviations in simulated precipitation can only be attributed to soil moisture impacts if the systematic effects of soil moisture modifications are larger than the inherent simulation uncertainty at the convection resolving scale.

We performed seven experiments with either modified soil moisture amount or distribution to address the effect of soil moisture on precipitation. Each of the experiments consists of ten ensemble members using the deep convection resolving COSMO model with a grid spacing of 2.8 km. In three out of the seven experiments, precipitation changes exceed the model spread in either amplitude, location or structure. These changes are caused by a 50% soil moisture increase in either the whole or part of the model domain or drying of the whole model domain. Increased or reduced soil moisture predominantly results in reduced precipitation rates. Replacing the soil moisture by a realistic field from different days has an insignificant influence on precipitation. We point out the need for uncertainty estimates in soil moisture studies based on convection resolving models.

1 Introduction

Convective precipitation changes rapidly in time and is very variable in space (Pedersen et al., 2010). The heterogeneity of convective precipitation and the interaction of different scales challenge atmospheric models on the global and regional scale. Nowadays, regional climate models operate with a horizontal resolution of 1 km and can represent convective processes explicitly to improve weather forecast (Mass et al., 2002). Nevertheless, precipitation formation results from a complex chain of atmospheric processes, which range from the microscale to the synoptic scale (Richard et al., 2007). Because many of these

processes remain unresolved, precipitation features a highly uncertain quantity.

Soil moisture determines the partitioning of energy and the land surface and controls how much energy is used to heat up the surface or to moisten the atmosphere. The surface temperature plays a crucial role in the initiation of convection, whereas the specific water content in the boundary layer modifies moist conditional instability. On the one hand, high surface temperatures initiate convection with a low soil moisture content. On the other hand, high soil moisture can destabilize the atmosphere by introducing water vapor in the lower troposphere resulting in an enhanced possibility for convection. Accordingly, there is no distinct effect from soil moistening or drying on precipitation intensification. In contrast, there exists a strong systematic influence of soil moisture changes on latent and sensible heat fluxes as well as on equivalent potential temperature, lifting condensation level, and convective energy (Barthlott et al., 2011). Despite these systematic effects, precipitation reacts less systematically to soil moisture variations (Barthlott and Kalthoff, 2011; Hohenegger et al., 2009). The distribution and inhomogeneity of soil moisture patterns may even initiate secondary circulation (Clark et al., 2004; Adler et al., 2011; Kang and Bryan, 2011; Dixon et al., 2013; Maronga and Raasch, 2013; Froidevaux et al., 2014).

There is no clear agreement on soil moisture - precipitation interaction in the literature: Barthlott et al. (2011) found a strong influence on precipitation, with changes larger than 500% for a soil moisture variation of $\pm 25\%$ in regions with low mountain ranges, and changes of up to -75% for domains with higher mountain ranges. They could not identify significant differences between planetary boundary layer driven and synoptically forced conditions. Hauck et al. (2011) determined large systematic differences between simulated and observed soil moisture. The influence on simulated precipitation in their study was complex and strongly dependent on the chosen case and domain. A dependency of all convective indices on the equivalent potential temperature was found by Kalthoff et al. (2011) over different orographic terrains. However, convection was predominantly initiated over mountain crests, independently of the instability indices, but with smaller convective inhibition (CIN). The dependency of equivalent potential temperature on soil moisture was found to be influenced by surface inhomogeneity. Barthlott and Kalthoff (2011) provide a sensitivity study in which the soil moisture was changed by $\pm\,50\%$ in steps of 5%. While the study reveals a systematic effect on the 24 hours precipitation sum for reduced soil moisture, precipitation is not systematically modified by wetter soil moisture conditions.

Large variations in the results may partly be attributed to model uncertainty. Hohenegger and Schär (2007) investigated the error growth of random perturbation methods in cloud-resolving models using time shifted model simulations and perturbed temperature fields in the initial conditions. In their model study, using a model resolution of $2.2\,\mathrm{km}$, a rapid error growth was found far away from the perturbed regions, but growth of uncertainties is limited by the large-scale atmospheric environment. A further aspect causing model uncertainty is model resolution, especially regarding the influence on convection. Different results for soil moisture-precipitation feedback also appear in simulations with explicitly resolved and differently parametrized convection (Hohenegger et al., 2009). Hohenegger et al. (2008) found different results in sign and strength of the influence of soil moisture that depend on the model resolution. Simulations with explicitly resolved convection indicate a negative soil moisture-precipitation feedback that is in agreement with many other studies, summarised by Barthlott and Kalthoff (2011).

In numerical weather prediction models, soil moisture perturbations are used to generate the ensemble members. Weather services include soil moisture perturbation in data assimilation for their ensemble forecast systems. For example, MeteoSwiss

uses the method from Schraff et al. (2016) in the COSMO model to achieve a model spread, especially in summer (pers. com. Daniel Leuenberger, MeteoSwiss). The evaluation of the AROME-EPS (MeteoFrance) ensemble prediction system from Bouttier et al. (2016), which also includes soil moisture perturbations, shows a lack of spread in forecasted precipitation rates. This raises the question if soil moisture perturbations can cause sufficient differences in simulated precipitation. This question

5 will be addressed in the present study. As Richard et al. (2007) already stated, convective precipitation output strongly depends on the model setup, such as the prescribed initial conditions and boundary data. In the present study, we provide a description of changes in simulated precipitation resulting from different amounts or changed pattern of soil moisture together with an assessment of the uncertainty in precipitation caused by random processes in the model. The uncertainty is estimated from an ensemble generated with different boundary conditions by slightly shifting the model domain. Based on a large number of sim-

10 ulations with slightly changed model setup, the systematic influence of different soil moisture modifications on precipitation can be identified and quantified.

Model simulations are conducted with the regional model COSMO (section 2.1). The soil moisture experiments and the ensemble approach are presented in sections 2.2 and 2.3, respectively, for a single case with convective precipitation. An overview of the synoptic conditions for this convective case is provided in section 3. An estimate on the model uncertainty based on

15 the CTRL-ensemble is calculated in section 4.2. With the given uncertainty range, the significance of changes in precipitation caused by soil moisture changes compared to the model spread is assessed in section 4.3, and systematics in the soil moisture impact are investigated in section 4.4.

2   Modelling approach

20 ### 2.1   Numerical setup

We performed 24-hour integrations to simulate the convectively induced precipitation on 3 August 2012 over the area around Hamburg (Germany), using the non-hydrostatic model COSMO (Version 4.22, Schättler et al., 2009) with a horizontal resolution of $0.1°$ ($≈1$ km). The chosen domain covers 400×450 grid points over Northern Germany (Fig. 1). 50 vertical hybrid Gal-Chen levels range from the surface up to a height of 22 km. The lowest level has a vertical resolution of 20 m. Boundary

25 and initial conditions are provided by the COSMO operational analysis with a resolution of 2.8 km.

The horizontal resolution of approximately 1 km allows for an explicit representation of deep convection, and thus provides much more accurate simulations of convective precipitation than resolutions that require convection parametrizations (Leutwyler et al., 2016, and references therein). Shallow convection is parametrized using the Tiedke Scheme (Tiedtke, 1989). Land surface processes are calculated by the interactive soil and vegetation model TERRA-ML, and coupled to the atmospheric

30 module (Doms et al., 2011). The coupled soil model includes seven soil levels from the surface down to a depth of 14.58 m with the uppermost layer having a vertical extension of 5 mm.
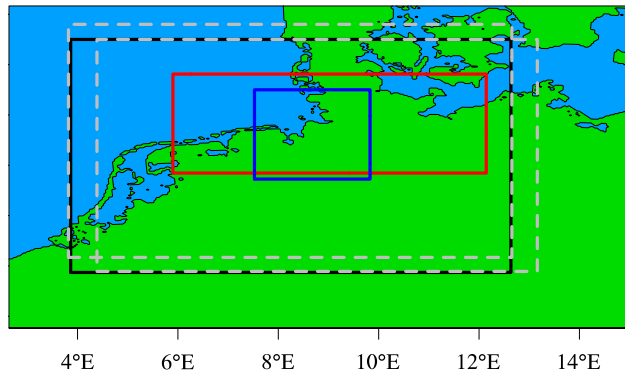
Figure 1. Model domain over Northern Germany given by the black rectangle for the CTRL run. Dashed gray rectangles describe the model domain shifted by 30 grid points to the north and to the east, respectively. The two analysis areas are marked with red and blue rectangles.

2.2 Soil moisture experiments

To address the potential effect of soil moisture amount and local distribution on precipitation, the soil moisture content provided in the initial conditions was modified (Table 1). Two ways of changes in the soil moisture field were applied: extreme artificial changes, which induce a large impact on precipitation and modifications in a physically feasible, realistic range (Fig. 2). As
5   an extreme modification, the total drying of the soil is implemented by setting the soil moisture content to zero (Fig. 2c). Soil moisture increase is achieved by an increase of 50% (Fig. 2d) in all soil layers. Those changes are first applied over all land points in the model domain (DRY_a and MOI_a, Table 1, Fig. 1) and second over land points in the domain framed in red in Fig. 1, hereafter referred to as area "red" (DRY_p and MOI_p, Table 1). Another artificial modification is achieved by a redistribution of the soil moisture into four alternating bands (BAND) with 50% increased and reduced soil moisture, respectively (Fig. 2b).
10   A large range of possible soil moisture effects is covered with these modifications. More realistic and less intense modifications are implemented by replacing the original soil moisture pattern by a real pattern at another day (Fig. 2e). For this purpose, soil moisture fields from 19 July 2012 and 20 August 2012 are used (Fig. 3). On the August 20, the soil moisture content in the uppermost soil layer (5 mm) is around 1.2 mm [H_2O] averaged over all land grid points in the model domain which is 0.3 mm lower than on the simulated day (3 August 2012). On 19 July 2012 soil moisture content was slightly below the 50%-artificial
15   increase (Fig. 3).
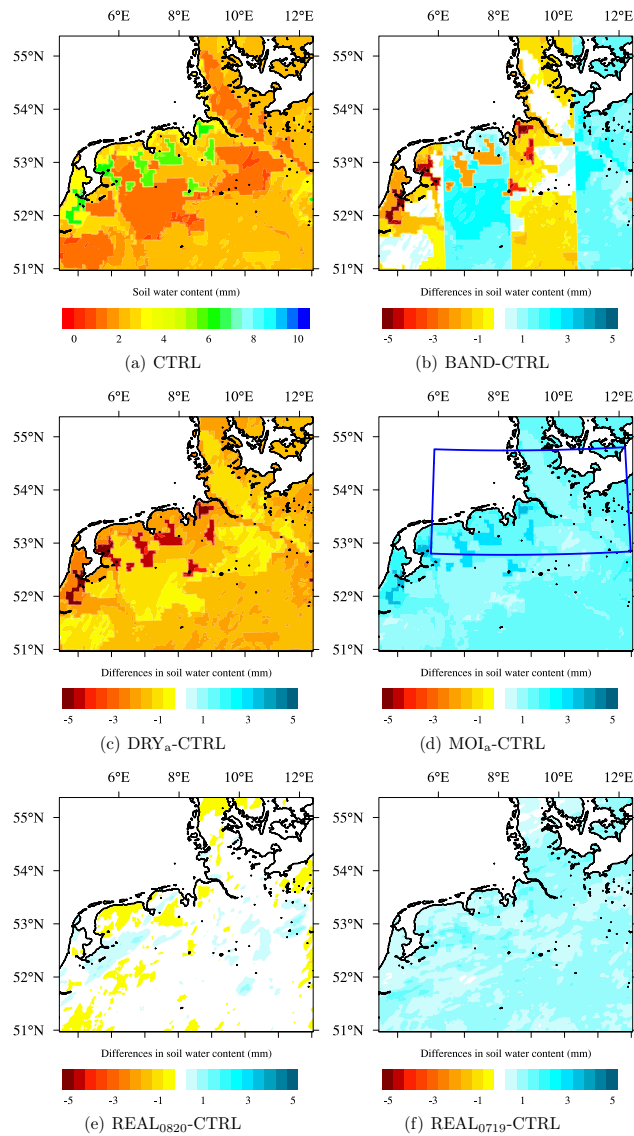
Figure 2. (a) Soil moisture for CTRL run and differences between CTRL run and (b) BAND run, (c) DRY_a run, (d) MOI_a run, (e) $REAL_{0820}$ run and (f) $REAL_{0719}$ run in the uppermost soil layer. Blue rectangle indicates the region where soil moisture was changed in DRY_p run and MOI_p run.

2.3 Ensemble approach

To quantify the relevance of the results from the soil moisture modifications, the model uncertainty and variability is estimated with a novel and simple approach. Perturbations are implemented in the simulations by shifting the domain boundaries by ten to 30 grid points north- and eastwards (Table 2, Fig. 1). These perturbations allow for estimating the uncertainty caused
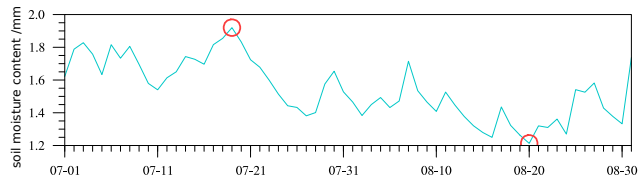
Figure 3. Time series for soil moisture content in the uppermost soil level averaged over the analysis domain "red". Red circles indicate the soil moisture values, which were used to perform simulations with soil moisture from another day.

Table 1. Model simulations with modified soil moisture (SM). Simulations are named by the applied soil moisture modification and with a for whole model domain and p for modification in a subdomain (partly). Simulations with additional random changes caused by spatial shifting of the domain are denoted with ii and jj, which represents number of grid points by which the model domain is shifted (For details see Table 2).

reference simulation for ensemble	Characteristics		ensemble generation	
	modification	area		
CTRL			CTRL-LOCii jj	TIMEtt
DRY$_a$	dry out	whole model domain	DRY$_a$-LOCii jj	
DRY$_p$	dry out	area "red"	DRY$_p$-LOCii jj	
MOI$_a$	50% increased SM	whole model domain	MOI$_a$-LOCii jj	
MOI$_p$	50% increased SM	area "red"	MOI$_p$-LOCii jj	
BAND	four bands	whole model domain	BAND-LOCii jj	
REAL$_{0820}$	SM from 20.08.12	whole model domain	REAL$_{0820}$-LOCii jj	
REAL$_{0719}$	SM from 19.07.12	whole model domain	REAL$_{0719}$-LOCii jj	

by the chaotic behavior of the atmospheric system and are superimposed on all systematic and physical changes caused by the soil moisture perturbations. This method conserves the structure of all meteorological input fields and does not create errors on a scale that can interact with the analysed processes e.g. by creating small-scale secondary circulations. Furthermore, shifting start times of the simulations (Hohenegger and Schär, 2007) provide an additional degree of uncertainty with the same advantages as the domain shift. A time shift of one to six hours is also applied to the CTRL run to allow for a fair comparison of uncertainty estimates. The ensemble, further called the CTRL-ensemble, comprises 17 independent model simulations including the reference simulation (CTRL-run) to estimate the uncertainty. This ensemble generating approach, including shifted model domain, is applied to every simulation with modified soil moisture pattern (Table 2).

Table 2. Uncertainty ensemble with randomly changed model simulations by model domain shifting (LOC) and the number of shifted grid points, and caused by shifting the model start time (TIME). The time shift is given in hours. The lower left (LL) corner of the simulation domains is given in geographical (rotated) coordinates with the north pole being shifted to 40° N and -170° E.

run	LL corner in ° N	LL corner in ° E	starttime (UTC)
CTRL	50.87 (1.0)	15.55 (3.5)	0:00
LOC 00 10	50.97 (1.1)	15.56 (3.5)	0:00
LOC 00 20	51.07 (1.2)	15.57 (3.5)	0:00
LOC 00 30	51.17 (1.3)	15.59 (3.5)	0:00
LOC 10 00	50.88 (1.0)	15.39 (3.4)	0:00
LOC 10 10	50.98 (1.1)	15.40 (3.4)	0:00
LOC 10 20	51.08 (1.2)	15.42 (3.4)	0:00
LOC 20 00	50.89 (1.0)	15.23 (3.3)	0:00
LOC 20 10	50.98 (1.1)	15.25 (3.3)	0:00
LOC 20 20	51.08 (1.2)	15.26 (3.3)	0:00
LOC 30 00	50.89 (1.0)	15.08 (3.2)	0:00
TIME 01	50.87 (1.0)	15.55 (3.5)	1:00
TIME 02	50.87 (1.0)	15.55 (3.5)	2:00
TIME 03	50.87 (1.0)	15.55 (3.5)	3:00
TIME 04	50.87 (1.0)	15.55 (3.5)	4:00
TIME 05	50.87 (1.0)	15.55 (3.5)	5:00
TIME 06	50.87 (1.0)	15.55 (3.5)	6:00

3 Convective case study and the effect of soil moisture

The chosen convective case of 3 August 2012 is characterized by a low pressure system in the northern Atlantic, west of Great Britain (Fig. 4). The associated cold front moved across Germany and resulted in heavy precipitation over Poland where air masses converged. Next to the major precipitation events in the east, another local precipitation cell developed close to
5 Hamburg where a slight enhancement in convective available potential energy (CAPE) values and high clouds occurred (Fig. 4). This strong local precipitation cell was detected by the rain radars over Hamburg at 14:11 UTC with rain rates between 10 to $100\,\mathrm{mm\,h^{-1}}$ (Fig. 5). The COSMO simulations showed maximum values of $12\,\mathrm{mm\,h^{-1}}$ between 13:00 and 17:00 UTC. The simulated precipitation onset is around 10:00 UTC (Fig. 7a). Before precipitation onset, high CAPE-value confirms the convective nature of the precipitation (Fig. 7b). High CAPE-values enable the development of strong convective precipitation,
10 when CIN can be exceeded.
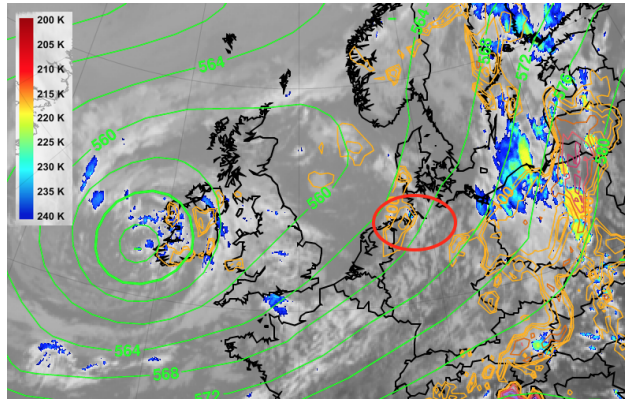
Figure 4. EUMEtrain infra-red satellite image of Europe on 3 August 2012 at 12:00 UTC. Colors are cloud top temperatures showing high clouds. Green contours are geopotential height at 500 hPa and orange contour lines are CAPE-values provided from ECMWF NWP (www.eumetrain.org). The area of interest is marked with a red circle. Note the CAPE values within the marked area and the high clouds, which correspond to the intense precipitation.
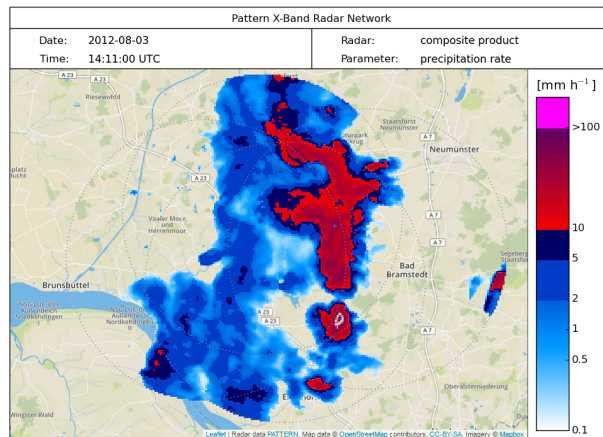
Figure 5. Radar composite from high resolution radars (Lengfeld et al., 2014) showing the precipitation rate over Itzehoe on 3 August 2012 at 14:11 UTC.

4 Results

4.1 Soil moisture influence on convection-related variables

While the passing front is the main mechanism for the lifting of air masses in the performed simulations, soil moisture is important for the stability of the atmosphere and thus affects precipitation initially triggered by the synoptic system. In complete dry

5 conditions (DRY_a and DRY_p), the latent heat flux is zero and the sensible heat flux needs to balance alone the net radiation flux and the soil heat flux (Fig. 6). Without latent cooling, the temperatures at $2\,m$ altitude in the simulation with dry soil conditions (DRY_a and DRY_p) are the highest, whereas dew point temperatures are the lowest (Fig. 6c and d). With more humidity in the atmosphere, less adiabatic cooling due to lifting is required for condensation. The resulting shift in the condensation level to lower altitudes can reduce CIN and increase CAPE. CIN is only reduced in the first hours of the simulation before solar

10 radiation heats the surface (Fig. 7). When the surface heats up in simulations with low soil moisture content (DRY_a and DRY_p), CIN is continuously lower than in simulations with higher soil moisture content (MOI_a and MOI_p). The further development of CIN is strongly affected by the feedback from precipitation.

All mentioned quantities react systematically to changes in soil moisture amount (Fig. 6). Convective precipitation is more likely with reduced CIN and its strength depends on CAPE. However, precipitation does not respond to changes in soil mois-

15 ture. The main reason for this systematic behavior mentioned in literature is the dependency of CIN on the soil moisture (Kalthoff et al., 2011). Even though the changes in precipitation caused by changing soil moisture are more complex, precipitation is certainly influenced by the soil moisture. However, the changes in precipitation caused by soil moisture modifications are often less significant than changes caused by synoptic forcing which will be demonstrated in the following.

4.2 Estimate of model uncertainties

20 An estimate of the model uncertainty is determined from the CTRL-ensemble. The assessment of this uncertainty is done statistically by using the SAL score (Wernli et al., 2008), which assigns values for differences in structure, S, amplitude, A and location, L, between precipitation patterns at every single output-timestep (15min). These three parameters of the SAL score are shortly introduced in the following.

Amplitude, A, describes the differences of precipitation amount over the whole analyzed domain:

25
$$A = \frac{D(R_{\mathrm{dif}}) - D(R_{\mathrm{ref}})}{0.5[D(R_{\mathrm{dif}}) + D(R_{\mathrm{ref}})]}. \tag{1}$$

The precipitation amount averaged over the whole domain $D(R)$ is:

$$D(R) = \frac{1}{N_{\mathrm{GP}}} \sum_{(i,j) \in \epsilon} R_{ij}, \tag{2}$$

where R_{ij} is the precipitation rate at the grid point with indices i, j, and N_{GP} the number of all grid points in the analyzed domain. $D(R_{\mathrm{dif}})$ denotes the averaged precipitation amount for shifted simulations and $D(R_{\mathrm{ref}})$ the reference simulations (not

30 shifted) in the CTRL-ensemble.
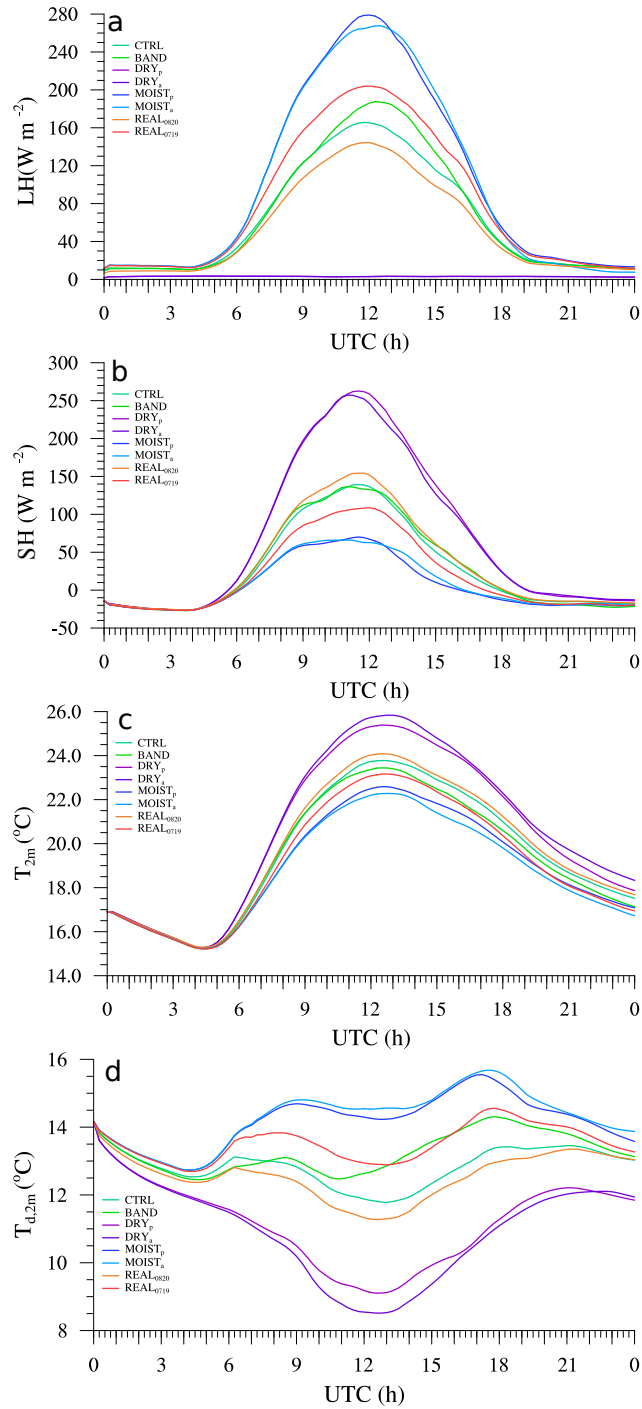
Figure 6. Timeseries of (a) latent heat fluxes, LH, (b) sensible heat fluxes, SH, (c) 2m-temperature and (d) dew point temperature, T_d in the reference simulations of the ensembles with different soil moisture modifications for 3 August 2013. Averaged over area "red".
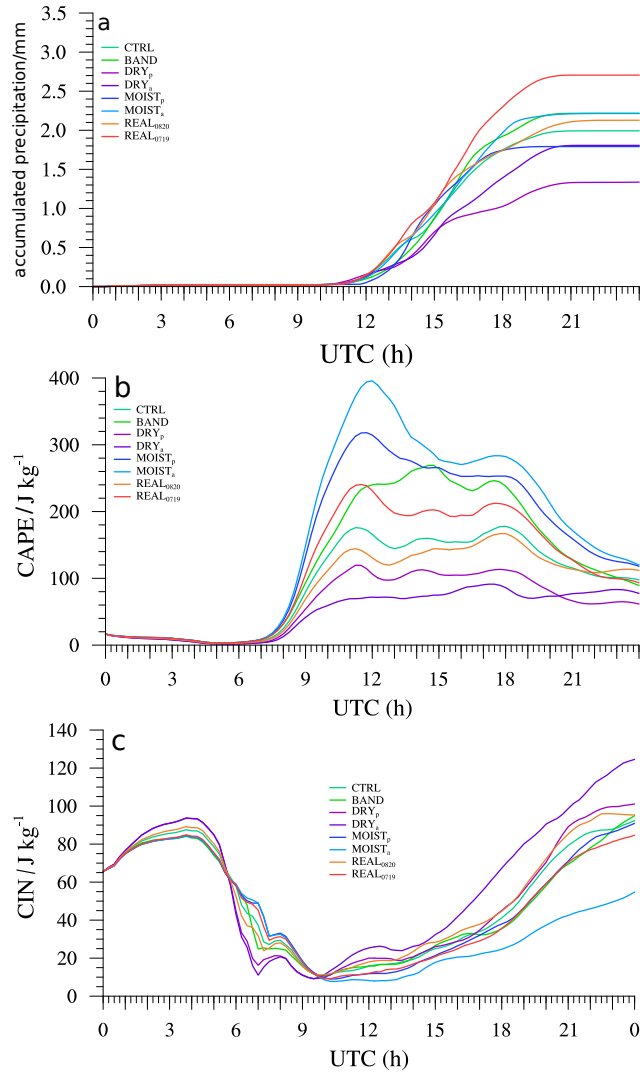
Figure 7. Timeseries of (a) accumulated precipitation, (b) CAPE and (c) CIN for different model simulations for 3 August 2013. Averaged over "red" area.

The location parameter, L, compares the location of precipitation in two model simulations in two steps. First, the normalized distance of the centers of mass, $\mathbf{x}(R)$, of the precipitation patterns in each model simulation is calculated:

$$L_1 = \frac{|\mathbf{x}(R_{\text{dif}}) - \mathbf{x}(R_{\text{ref}})|}{d}, \tag{3}$$

where d denotes the maximum possible distance within the analyzed domain. Secondly the distances from the center of mass of all M individual cells, $\mathbf{x}_n$, to the center of mass for the whole precipitation field, $\mathbf{x}$, are calculated as:

$$r(R) = \frac{\sum_{n=1}^{M} R_n |\mathbf{x} - \mathbf{x}_n|}{\sum_{n=1}^{M} R_n} \tag{4}$$

The distances resulting from the reference simulation and shifted simulations are then compared:

$$L_2 = 2 \left[\frac{|r(R_{\text{dif}}) - r(R_{\text{ref}})|}{d} \right]. \tag{5}$$

Both components of L are added.

Structure component, S, gives a hint whether the precipitation patterns tend to more convective precipitation with small but more peaked rain objects or shallow precipitation with larger objects, but weaker precipitation intensity. A volume, $V(R)$, is calculated by dividing the precipitation sum for a cell n, $\mathcal{R}_{\text{ij}}$, calculated over the ϵ grid cells of n, by the maximum precipitation of this cell, $R_n^{\max}$:

$$V_n = \sum_{(\text{i,j}) \in \epsilon} \frac{\mathcal{R}_{\text{ij}}}{R_n^{\max}}, \tag{6}$$

$$V(R) = \frac{\sum_{n=1}^{M} R_n V_n}{\sum_{n=1}^{M} R_n}. \tag{7}$$

With the volume, $V(R)$, over all precipitation cells M the structure component can be calculated similarly to Eq. (1):

$$S = \frac{V(R_{\text{dif}}) - V(R_{\text{ref}})}{0.5[V(R_{\text{dif}}) + V(R_{\text{ref}})]}. \tag{8}$$

For more detailed information on SAL see Wernli et al. (2008).

The simulations are compared for the period from 10:00 UTC to 18:00 UTC, covering the precipitation event. Those with a shifted model start or domain are compared to the CTRL run. Amplitudes, A, are positive or negative dependent on which run is used as reference (ref) or as a comparison simulation (dif). To avoid an uncertainty range shifted toward one sign, all comparisons are additionally performed after swapping ref and dif run, providing a symmetric uncertainty distribution. The uncertainty estimate (Fig. 8) encompasses a sample of 122 (permutation of simulation couples) $\times$ 32 (time steps) values, although not all of them are independent of each other.

As parameters A and S are correlated (Fig. 8), a reduction in precipitation amplitude is related to too small and/or peaked precipitation objects, whereas an increase in precipitation amount goes along with larger and/or shallower rain objects. Largest amplitude deviations between the different runs arise in the first hours of the analysis time from 10:00 to 11:30 UTC (Fig.
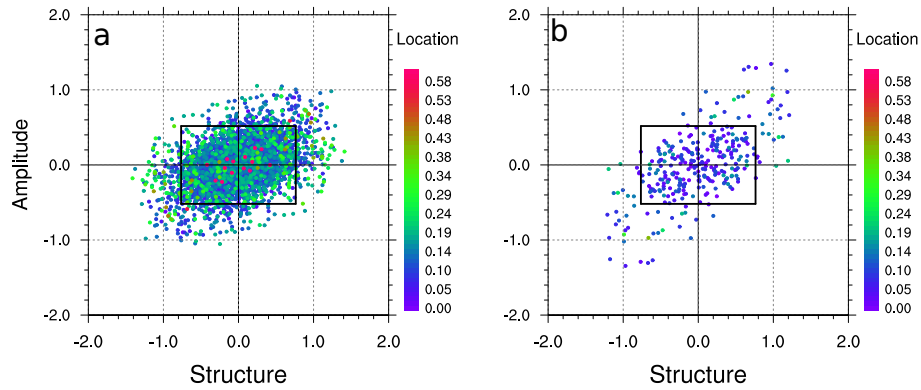
Figure 8. SAL results for the CTRL-uncertainty-ensemble between 10:00 UTC and 18:00 UTC separated for uncertainty generated by the spatially shifted model domain (a) and by the delayed model start (b). Structure is represented on the x-axis, amplitude on y-axis and location by marker colours. Every marker shows a comparison between two model simulations at a single time step. Simulations with shifted model domain are represented by filled dots and simulations with shifted model start time by rectangles. The grey rectangle delimits the region between the 5% and 95% percentiles in structure and location amplitude.

9). This coincides with the time of the onset of the precipitation event which differs in the different simulations (Fig. 7) and therefore causes the largest uncertainties. The end of the precipitation event is not considered in the chosen time range.

A large shift in model start time leads to higher uncertainties (Fig. 9). Changes due to the spatially shifted model domain do not depend on the distance of the boundary shift. The deviations from CTRL for simulations with shifted boundaries are not

5 caused by a direct change of physical parameters such as temperature, but the differences emerge because the synoptic forcing at the lateral model boundary differs. Deviations are further caused by changes in the lower boundaries, such as changing areas of sea- or landcover. For example, a model domain with north- or westward shifted boundaries includes more grid points over sea surface. The simulation with the strongest westward shift (LOC3000) shows the largest changes in precipitation amplitude (Fig. 9). However, the strongest shift (30 km north in LOC3000) affects the precipitation amplitude less strongly than a smaller

10 shift (20 km north in LOC2000), which includes a smaller fraction of sea surface (Fig. 9).

To address the dependency of SAL score on the chosen analysis area, two different analysis areas are chosen (Fig. 1). The area framed in blue in Fig. 1 includes mainly small convective cells and the area framed in red includes the whole precipitation field. For those two analysis areas, two simulations are compared. We define the model uncertainty for this study as the range between the 5% and 95% percentile for structure and amplitude and up to the 90% percentile for location. According to this

15 definition, the uncertainty range is ±0.77 (±0.86) in structure, ±0.54 (±0.69) in amplitude, and up to 0.20 (0.29) in location for analysis area "red" ("blue"). Changes are defined as significant when the response to the soil moisture modification is larger than the generated background noise from spatially shifted model domains or delayed model start time. A residual probability of 10% remains that the changes do not result from soil moisture modification. When comparing the CTRL run (Fig. 10a) and the simulation with shifted boundaries (Fig. 10b), differences in the single cells in the western part of the domain, partly

20 over the North Sea, and in the structure of the large precipitation pattern in the eastern part of the domain become obvious.
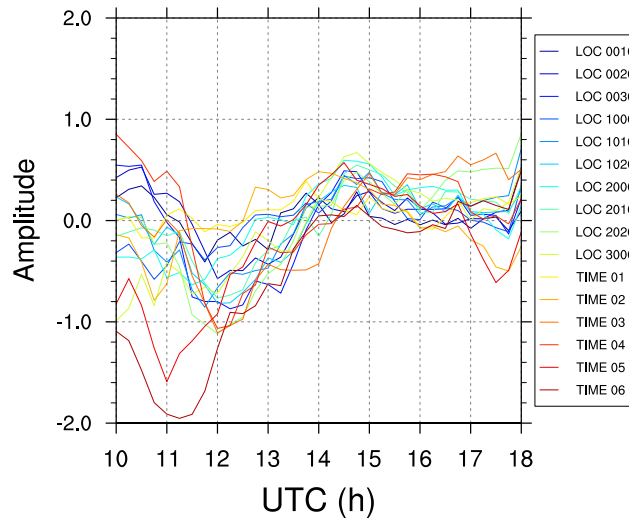
Figure 9. Amplitude values from Fig. 8 for comparisons to the CTRL run only, for single time steps.

These differences are caused by shifting the boundary domain by ten grid points (10 km). The extreme modifications in soil moisture cause even more obvious differences in the precipitation patterns. The increase of soil moisture in either the whole domain or in a sub domain dramatically changes the location of the precipitation cells in the analysed time steps (Fig. 10e and f). In the moist simulations (MOI_a and MOI_p), the strongest precipitation occures north-east of the Elbe-estuary. This region is mainly free from precipitation in the dry simulations (DRY_a and DRY_p), and the precipitation occurs even further north-east of this particular region in the CTRL simulations (CTRL and LOC1000). Moderate changes in soil moisture, e.g. when applying realistic moisture fields of a different day, result in smaller changes in precipitation. The general pattern observed in the CTRL run remains the same in $REAL_{0820}$ and $REAL_{0719}$ (Fig. 10f and g).

4.3 Significant effects of soil moisture modification on precipitation

The huge number of model simulations (a complete ensemble for each soil moisture modification) and the uncertainty estimate from section 4.2 accounts for a quantitative evaluation of the significance of soil moisture influence on precipitation. Each ensemble with modified soil moisture is compared to the CTRL uncertainty ensemble by comparing ensemble members with same spatial model domain shift for every output-step. Within every ensemble the values are divided into those that exceed the uncertainty range given by blue rectangle in Fig. 11 derived from the uncertainty-ensemble, and those that lay within. The percentage of values exceeding the uncertainty range is calculated to decide whether a soil moisture modification leads to significant changes in precipitation (Table 3). Changes caused by a soil moisture modification are considered as significant if more than 10% of the values exceed the uncertainty range. The threshold is set to 10% because the uncertainty range is determined by considering the range between the 5% and 95% percentiles, which leaves 10% probability that the exceeding

value is still caused by model uncertainty.

The change in structure of precipitation caused by the soil moisture modification in the DRY$_{\mathrm{p}}$-ensemble exceeds the uncertainty range in only 5% of all cases (Fig. 11a and Table 3). For both scores, S and A, the percentage of exceeding values lies beneath the 10% threshold. Therefore, precipitation does not respond significantly to DRY$_{\mathrm{p}}$ modifications, but for the
5   parameter L in area "red" (Table 3). In contrast, the soil moisture reduction in the whole domain (DRY$_{\mathrm{a}}$-ensemble) affects the precipitation significantly (Fig. 11b). More than 50% of amplitude values exceed the uncertainty range, in some cases with values for A down to -1.8. For S, only 11% of the values exceed the uncertainty range. Nevertheless, this is enough to be regarded as a significant impact. The soil moisture increase in a sub domain only (MOI$_{\mathrm{p}}$-ensemble) results in significant changes in precipitation (Fig. 11c). As already seen in the DRY$_{\mathrm{a}}$-ensemble, the modification over the whole domain results in
10   a stronger precipitation response.

The redistribution of soil moisture (BAND-ensemble) does not lead to any significant effect (Fig. 11e), except for L in area "blue". It changes the heterogeneity of the soil moisture by reducing small-scale structures, but induces stronger variations on a large scale. Thus, secondary circulations can develop on a different scale. This is in accordance with Adler et al. (2011) and Kang and Bryan (2011) who both found an influence of the redistribution of soil moisture on the location of convective
15   initiation. Therefore, area "blue", mainly containing small convective cells, is influenced to a greater extent than area "red", which has a large advected precipitation band.

Even slight modifications of soil moisture, as Klüpfel et al. (2011) performed by using different initialization of soil moisture, lead to different precipitation patterns. In the present study, using soil moisture from another day also changes precipitation. But those changes do not exceed the model uncertainty in more than 10% of all values in the present case. Accordingly, physically
20   realistic changes in soil moisture lead to changes in precipitation not larger than changes that can also be caused by choosing a slightly different model setup.

4.4   Systematic behavior to soil moisture changes

Having determined the significance of the strength of changes in precipitation, this section handles the systematics of changes. Significant changes do not necessarily imply systematic changes. Changes in L are not analyzed as this quantity provides
25   no direction. While predominantly negative amplitude changes occur in DRY$_{\mathrm{a}}$ (Fig. 11b), changes in MOI$_{\mathrm{p}}$ (Fig. 11c) are significant, but random. Structure and amplitude change in both positive and negative directions. In none of the soil moisture experiments S and A are correlated (Fig. 11). To carve out any systematic effect, the averaged values of amplitude and structure are compared to the average of the uncertainty-ensemble. The sample for the SAL results for the uncertainty-ensemble is symmetric and therefore the average over all values is zero. A significant difference of the averaged values from zero hints at
30   the systematics. Whether the averaged values differ significantly from zero is tested statistically by:

$$\widehat{z}_{\mathrm{sys}} = \frac{\overline{x}_1 - \overline{x}_2 - E[\overline{x}_1 - \overline{x}_2]}{\sqrt{\widehat{\sigma}^2[\overline{x}_1 - \overline{x}_2]}}. \tag{9}$$

$\overline{x}_1$ and $\overline{x}_2$ denote the averaged values of S or A for the two compared simulations, $E[\overline{x}_1 - \overline{x}_2]$ is the expected value for the differences between the two simulations and is expected to be zero for the null hypothesis, $\widehat{\sigma}^2[\overline{x}_1 - \overline{x}_2]$ is the variance of the

Table 3. Percentages (p_S, p_A, p_L) of values S, A and L that exceed the model uncertainty by 90%. Uncertainty is in a range from $[-0.767, 0.767]([-0.857, 0.857])$ in structure, $[-0.538, 0.538]([-0.690, 0.690])$ in amplitude and $0.200(0.288)$ in location for analysis area "red" ("blue"). Bold values exceed model uncertainties in more than 10% of the cases. Averaged values and their deviations $(\overline{S} \pm \hat{\sigma}^2, \overline{A} \pm \hat{\sigma}^2)$. Bold values are those mean values that differ significantly from the mean of the uncertainty-ensemble after Eq. (9) for confidence interval of 90%.

Ensemble	Structure			Amplitude		Location
	p_S	$\overline{S} \pm \hat{\sigma}^2$	p_A	$\overline{A} \pm \hat{\sigma}^2$		p_L
			analysis area "red"			
DRY$_p$	5.79	0.02 ± 0.0034	3.58	-0.13 ± 0.0011		**25.34**
DRY$_a$	**23.14**	$\mathbf{0.30 \pm 0.0063}$	**22.31**	-0.26 ± 0.0023		**53.72**
MOI$_p$	**15.98**	-0.12 ± 0.0051	**18.73**	-0.05 ± 0.0042		8.26
MOI$_a$	9.92	-0.10 ± 0.0043	**23.42**	$\mathbf{-0.18 \pm 0.0043}$		**26.72**
BAND	3.03	-0.04 ± 0.0025	0.55	0.00 ± 0.0001		6.61
REAL$_{0820}$	3.31	-0.03 ± 0.0022	0.28	-0.02 ± 0.0006		0.55
REAL$_{0719}$	2.48	0.05 ± 0.0024	1.65	0.09 ± 0.0008		1.65
			analysis area "blue"			
DRY$_p$	4.85	-0.10 ± 0.0033	9.39	$\mathbf{-0.29 \pm 0.0091}$		5.76
DRY$_a$	**11.82**	0.08 ± 0.0058	**51.21**	$\mathbf{-0.60 \pm 0.0068}$		**30.61**
MOI$_p$	**14.85**	-0.19 ± 0.0053	**12.12**	0.00 ± 0.0039		**12.42**
MOI$_a$	**15.76**	$\mathbf{-0.27 \pm 0.0058}$	**19.70**	-0.28 ± 0.0044		**27.58**
BAND	7.27	-0.10 ± 0.0045	0.91	0.07 ± 0.0014		**21.52**
REAL$_{0820}$	0.91	-0.04 ± 0.0026	0.30	-0.03 ± 0.0007		1.21
REAL$_{0719}$	1.21	-0.01 ± 0.0026	0.30	0.07 ± 0.0010		1.21

averages.

The average of the structure values differs significantly from zero only for the ensembles DRY$_a$ and MOI$_a$ in the analysis domain "red" and "blue", respectively (Table 3). The amplitude values frequently feature a systematic response in both domains: Modifications of soil moisture, increases as well as decreases, result in in reduced precipitation rates. This implies a negative and positive feedback, respectively. The positive feedback activated by reduced soil moisture is in consent with a case study by Barthlott and Kalthoff (2011). However, increased soil moisture amount can lead to an increase or decrease in precipitation, dependent on the strength of the increase. In contrast, Cheng and Cotton (2004); Ek and Holtslag (2004); Martin and Xue (2006); Hohenegger et al. (2009); Weverberg et al. (2010) found a negative feedback in convection resolving simulations.

The strength of the deviation depends on the strength of the modification. While a partial increase in soil moisture does not lead to systematic changes, an overall enhancement has a systematic effect. The effect of dry soil exceeds the effect of soil moisture enhancement and shows systematic effects for both implementations (drying the entire domain or only parts). The

effects are stronger for overall modifications. Comparing the results for both regions, the averaged differences calculated for region "blue" exceed those of region "red" because convective cells are more influenced by soil moisture changes.

5 Conclusions

In the present case study we conducted seven ensembles for different perturbations in soil moisture amount and pattern. Each ensemble encompasses ten variations in the model domain. The soil moisture perturbations include strong artificial changes by drying and wetting the model domain and realistic changes implemented by replacing the initial soil moisture field with real soil moisture patterns from another day. The ensembles are generated by conducting the simulation over slightly different model domains, created by shifting the domain location by 10 to 30 km to change the boundary conditions. These changes cause a large spread in the convective case because the day (3 August 2012) was synoptically forced by a low pressure system over the Atlantic. Conclusively, only two ensembles show significant changes caused by modified soil moisture amount in intensity, local distribution and amount of convective precipitation which are assessed with the SAL-score. The amplitude, determining the difference in the amount of precipitation, is most often systematically reduced within an ensemble. No overall systematic was found because wetting and drying of soil can both result in reduced precipitation amount. The structure, which describes the spatial variability of the precipitation field, can be either enhanced or decreased at different times and in different ensemble members. This might be explained by a delayed onset of precipitation caused by soil moisture modifications. A local displacement in the precipitation cells is found for three and four out of five artificially changed soil moisture patterns in the two analysis regions. The changes in precipitation for the simulations with realistic soil moisture patterns are not significant. Limitations in the present study are: 1) the restriction to a single case study. Thus, no generally valid results can be found. However, this study presents a proof of concept and should be further conducted with cases, that are less affected by frontal systems. In those cases, a stronger influence of soil moisture on precipitation properties can be expected. 2) The dependency of the results on the chosen analysis area, which in turn shows the complexity of the results. 3) The dependency of the uncertainty estimate on synoptic forcing and size of the model domain. The ensemble spread might become smaller with weaker synoptic forcing and with a larger model domain. However, a smaller model spread would strengthen the importance of soil moisture influence as it is expected in these cases.

In summary, we could prove the concept that by shifting the model domain sufficiently large model spread can be achieved without being influenced by generated patterns. Such an estimate of the model spread is necessary in soil moisture studies to separate the response to soil moisture changes from inherent forecast uncertainty at deep convection resolving grid spacing. We further showed that in synoptically driven situations the effect of soil moisture remains uncertain and further investigation is necessary.

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Data availability

Model output data are stored at DKRZ computers and can be provided on demand.

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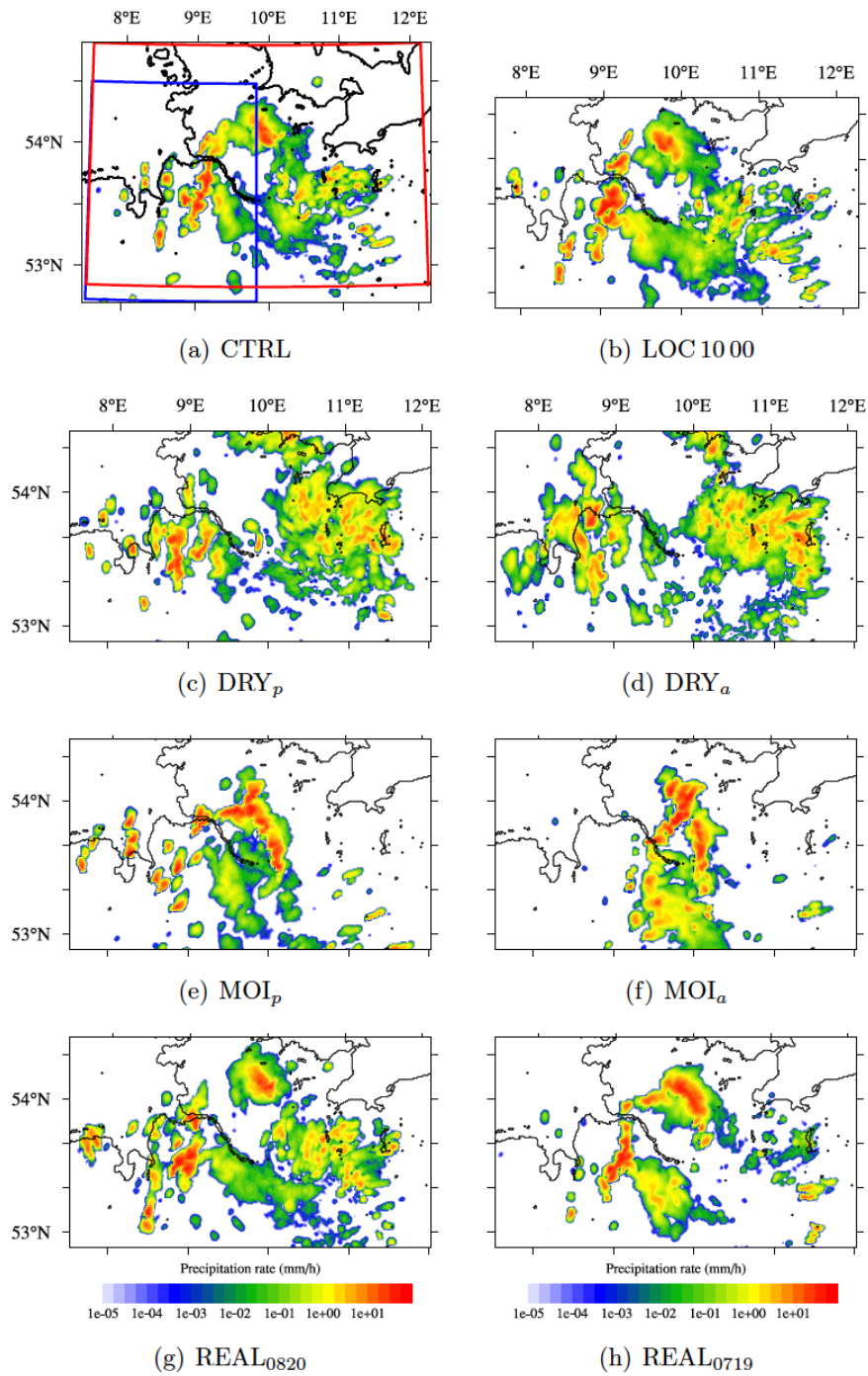
Figure 10. Precipitation rate at 14:45 UTC for (a) CTRL run , (b) LOC 10 00 and (c-h) different soil moisture modified simulations.
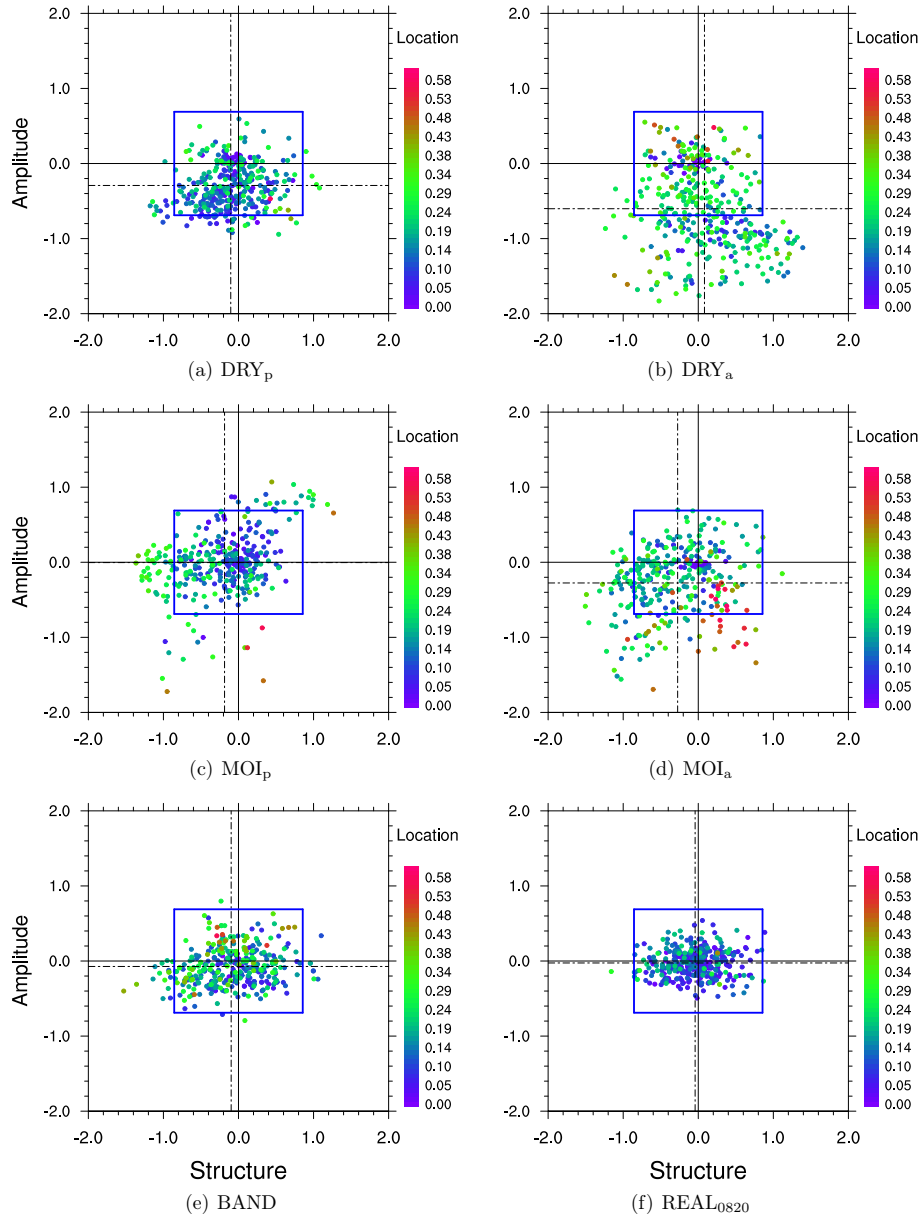
Figure 11. SAL scatter plot: Comparison of ensembles (a) DRY_p, (b) DRY_a, (c) MOI_p, (d) MOI_a, (e) BAND and (f) MOI_{0820} with the uncertainty-ensemble for area "blue". Dashed lines represent the averages.