

Are precipitation anomalies associated with aerosol variations over Eastern China?

Xiangde Xu¹, Xueliang Guo^{1,2,*}, Tianliang Zhao^{3,*}, Xingqin An¹, Yang Zhao¹, Jiannong Quan⁴, Fei Mao¹, Yang Gao⁴, Xinghong Cheng¹, Wenhui Zhu¹, and Yinjun Wang¹

¹State Key Laboratory of Severe Weather (LASW), Chinese Academy of Meteorological Sciences, Beijing 100081, China,

²Key Laboratory for Cloud Physics, Chinese Academy of Meteorological Sciences, Beijing 100081, China, ³Collaborative

³Innovation Center on Forecast and Evaluation of Meteorological Disasters, Key Laboratory for Aerosol-Cloud-Precipitation of China Meteorological Administration, Nanjing University of Information Science & Technology, Nanjing 210044, China,

⁴Beijing Weather Modification Office, Beijing 100089, China

* Correspondence to: Tianliang Zhao (tlzhao@nuist.edu.cn); Xueliang Guo (guoxl@camsma.cn)

Abstracts. In Eastern China (EC), the strong anthropogenic emissions deteriorate atmospheric environment, building a south-north zonal distribution of high aerosols harbored by the upstream Tibetan and Loess Plateaus in China. This study climatologically analyzed the interannual variability of precipitations with different intensities in association with aerosol variations over the EC region from 1961 to 2010 by using precipitation and visibility data in of more than 50 years and aircraft and surface aerosol data in recent years in China, and the impacts of aerosol variations on interannual variability of various precipitation intensities of precipitation events and their physical causes are investigated. We found that the frequency of light rain significantly decreased and the occurrence of rainstorm, especially the extraordinary rainstorm significantly increased over recent decades. The extreme precipitation events presented the similar interannual variability pattern with the frequent haze events over the EC. Accompanied with the frequent haze events in EC, the light rain frequency significantly decreased and the extremely heavy precipitation event have occurred more frequently. During the 1980s, the regional precipitation trends in EC showed an obvious “transform” from more light rain to more extreme rainstorms. The running correlation analysis of interdecadal variation further verified that the correlation between the increasing aerosols and frequencies of abnormal precipitation events tended to be more significant in the EC. The correlation between atmospheric visibility and low cloud amounts, which are both closely related with aerosol concentrations, had a spatial distribution of “northern positive and southern negative” pattern, and the spatial distribution of the variability of regional rainstorm frequency was “southern positive and northern negative”. After the 1990s, the visibility in summer season deteriorated more remarkably, and light rain frequency decreased obviously while the rainstorm and extraordinary heavy rainfall occurred more frequently. There were significant differences in the interdecadal variation trends in light rain

32 and rainstorm events between the high aerosol polluted area in the EC and the relatively “clean area” in the western plateaus
33 of China. The aircraft measurements over the EC confirmed that the diameters of cloud droplets decreased under high
34 aerosol concentration condition, thereby inhibiting weak precipitation process.

35 **1. Introduction**

36 Under the background of global warming, the regional precipitation tends to have more complex temporal and spatial
37 distribution patterns. The variations of precipitation could be reflected by the different-grade precipitation, and even by
38 frequency changes of extreme precipitation events (Lau and Wu, 2007). Precipitation is not only influenced by
39 atmospheric circulation related with land-sea discrepancy and land-sea water vapor exchange, but also by local cloud
40 microphysical processes. Atmospheric aerosols might add cloud droplets number concentrations, and change cloud lifetime,
41 and modify precipitation (Khain et al., 2005; Rosenfeld et al., 2007; Rosenfeld and Coauthors, 2008; Stevens and Feingold,
42 2009; Fan et al., 2013). Aerosols might also change Asian monsoon system (Bollasina et al., 2011). The interaction of
43 aerosols and cloud-precipitation is still an important issue with large uncertainties for climate change (IPCC, 2013).

44 Since the middle 1980s, China has been experiencing a rapid development in industry and agriculture. As a result, a huge
45 amount of anthropogenic emissions and biomass burning increasingly released particulate matters into the atmosphere. There
46 was no obvious change in annual precipitation in China, but the extremely heavy rainfall area, mainly in the EC, had
47 expanded (Zhai et al., 1999). However, the regional annual precipitation, summer precipitation, and extreme precipitation
48 events had obvious rising tendencies in middle and lower Yangtze River Basin of EC (Wang and Zhou, 2005).

49 The numerical simulations also presented that the increase of aerosols could decrease the summer convective precipitation in
50 the intensity under 30 mm h^{-1} , and increase summer strong convective precipitation in the rates above 30 mm h^{-1} in China
51 (Guo et al., 2014). With a rapid increase of aerosols, not only light rain over wide areas could decrease, but also local
52 extremely heavy rain could be triggered, inducing frequent flood (Guo et al., 2014; Fan et al., 2015). Light rain tended to
53 decrease and at the same time the extremely heavy precipitation had increasing tendency in the EC (Choi et al., 2008; Qian
54 et al., 2007; Qian et al., 2009). This phenomenon might be the strong signal of climate variability connecting to global
55 warming together with the increased emissions of anthropogenic aerosols.

56 The previous investigations of this issue primarily focused on limited cases with large discrepancies (Rosenfeld et al., 2007;
57 Rosenfeld and Coauthors, 2008; Stevens and Feingold, 2009; Li et al., 2011; Fan et al., 2013). The climatic forcing of
58 aerosols on precipitation in a large-scale region and its physical causes has been poorly understood. The long-term visibility
59 data can be used to climatologically assess the air quality change (Wang et al. 2009; Che et al. 2009; Xia et al. 2006), as the
60 atmospheric visibility is a good indicator of air pollutant levels in the environmental atmosphere (Zhao et al., 2016). By
61 using precipitation and visibility data in a 50-year period and aircraft and surface aerosol observational data in recent years

62 in China, the climatic impacts of aerosols on interannual variability of various precipitation intensities and their physical
63 links were investigated in this study. In addition, the high aerosol concentrations are accumulated in the north-south direction
64 over EC in connection with the terrain effect of the Tibetan plateau and the Loess plateau in China (Xu, et al, 2016).The
65 polluted EC region and the clean plateaus in China may be the ideal places to identify the climate forcing of aerosols with
66 comparing the interannual variation trends in various precipitation intensities for exploring .the relation of precipitation
67 anomalies and aerosol variations.

68 2. Data

69 In this study, we classified extraordinary storm, large rainstorm, rainstorm, large rain, moderate rain and light rain
70 respectively with daily precipitation >200 and ranging between 100-200mm, 50-100mm, 25-50 mm, 10-25mm and
71 0.1-10mm in this study. The monthly data of precipitation events of extraordinary storm large rainstorm, rainstorm, large
72 rain, moderate rain, light rain from 601 stations in China over 1961-2010 were adopted from the National Meteorological
73 Information Center of China Meteorological Administration. In addition, the meteorological and environmental data
74 including haze days daily visibility and low cloud cover in 1961-2010 as well as the daily PM_{2.5} data of 946 stations in
75 2013-2014 in China were also used in this study.

76 In order to analyze the regional variations in aerosols over EC, we adopt the equivalent visibility by excluding the influence
77 of natural factors (Rosenfeld et al., 2007) on the observed visibility based on the meteorological data in 1961-2010 were
78 used in this study. The equivalent visibility was corrected VIS (dry) based on the following formula (1) under the relative
79 humidity from 40% to 90%.

$$80 \quad \frac{VIS}{VIS(dry)} = 0.26 + 0.4285\log(100 - RH) \quad (1)$$

81 The vertical changes of aerosol and cloud droplets size were comprehensively analyzed based on the aerosol-cloud data
82 observed from aircraft flights over Beijing and its surrounding regions during 2008-2010. The flight observed clouds were
83 mainly stratus cloud, stratocumulus and cumulus clouds, and the maximum detection altitude was 7000 m. There were 40
84 flights carried out in 2-6 h before precipitation. The flight area and tracks were shown in Fig. 1. The Passive Cavity Aerosol
85 Spectrometer Probe (PCASP-200, DMT Co.) was used for observing aerosol particle size in 0.1-0.3 μ m. The probe of Cloud,
86 Aerosol and Precipitation Spectrometer (CAPS, DMT Co.) was used for observing cloud droplets in 0.6-50 μ m. The probes
87 were returned to the DMT for standard calibration before starting measurements in each year. In addition, the probes were
88 calibrated using the spheres of polystyrene latex (PSL) of Duke Scientific Corporation for each month. Considering the
89 influence of cloud droplets on aerosol probing, the averaged aerosol concentration below 300m of cloud base was calculated
90 to represent aerosol concentration in clouds. The cloud droplet measurements were made within clouds at 100m height
91 intervals. The data were processed into two or more samples when the clouds were multiply layered.

92 **3 Haze distributions in Eastern China**

93 Due to the influence of the terrain on the typical westerly, the air flowing from the windward plateaus descended between
94 about 110°E and 125°E (upper panel of Fig. 2). Accompanying this strong downward current were weak winds in the
95 near-surface layers. The wind condition leads to accumulating air pollutants in EC. The weak wind and downward current
96 areas coincide well with the centers of frequent haze events in EC (lower panel of Fig. 2). The frequent haze pollution over
97 EC is associated with the “harbor” effect of the topography under specific meteorological conditions that trap air pollutants
98 (Xu et al., 2016). The EC is climatologically a region with frequent haze events over recent decades, where high aerosol
99 pollution could exert an impact on the regional variation in precipitation.

100 **4. Change trends in various precipitation intensities**

101 The trends interannual variation of precipitation events with various intensities of light rain, moderate rain, heavy rain,
102 rainstorm, the large rainstorm, extraordinary rainstorm over EC were comparatively analyzed in Figure 3. Regionally
103 averaged over EC, the trends in light rain frequency had significantly decreased, while the events of rainstorm including
104 large and extraordinary storm had increased significantly (Fig. 3a), although the moderate rain frequency trend slightly
105 declined, and the interannual change trend of large rain frequency was not significant (Fig. 3a). Especially since 1980s, the
106 extremely heavy precipitation events have become more frequent, showing frequent occurrences of disastrous rainstorm
107 along with the frequent haze in EC. Overall rainstorm extreme events were on the rise trend, but light rain tended to decline
108 significantly. In contrast, stations in the Tibetan Plateau (TP) with altitude >4000m, a relative clean area in China, were
109 selected for a statistical analysis of interannual variation trend of light rain frequency. The characteristic of the decreased
110 trend of light rain frequency was not significant in the TP over recent decades (Fig. 3b), which could imply aerosols
111 anomalies restraining light rain frequency over EC...

112 **5. Regional changes of precipitation events, haze and visibility**

113 We calculated the trends in interannual variations of precipitation and visibility at all the site in China (Fig. 4). The area with
114 the negative trends in light rain frequency quite well matched with the areas of negative trends in and positive trends in haze
115 frequency in EC (Figs. 4a, b and c). The light rain frequency reduction in China was closely associated with the
116 enhancement of aerosol levels in the atmosphere (Qian et al., 2009).

117 It is noteworthy that the negative trend areas of light rain covered the most sites in EC (Fig. 4c). This might be also closely
118 related with temporal and spatial variations of East Asian summer monsoon which offered a suitable dynamic background
119 for the effect of aerosols on clouds and precipitation. Figure 5a shows that a spatial distribution of the trends in rainstorm
120 frequency was “southern positive and northern negative” in summer during 1961-2010, while the correlations between

121 visibility and low-level cloud amount were distributed with the “northern positive and southern negative” pattern in EC
122 during 1961-2010 in EC (Fig.5b), indicating that the effect of aerosols on summer convective precipitation was more
123 obvious in southern part than that in northern part of EC.

124
125 There were obvious differences in the precipitation change rate of various precipitation intensities in the EC region (Fig.6a),
126 where the negative variability stations of light rain made up the majority (about 87.6%), the positive variability stations of
127 moderate rain were approximately equal to the negative ones (about 51%), the positive variability stations of large rain
128 (about 71.3%) were much more than the negative ones indicating the reverse trend. The positive variability stations of
129 rainstorm with daily precipitation >50mm, including catastrophic rainstorm over 100mm occupied obvious majority (about
130 78.9%). The increase of the anthropogenic aerosol particles in the atmosphere may suppress light rain (Qian et al., 2009),
131 and also enhance the rainstorm precipitation with more frequent events in EC. The anomalous changes of regional
132 precipitation from less light rain changed to more heavy rain and even the catastrophic rainstorm along with the frequent
133 haze pollution in EC.

134 Although precipitation events depended on dynamical and thermodynamic processes and water vapor source in the
135 atmosphere, aerosol’s “Albrecht effect” with increasing cloud droplet concentrations and decreasing cloud droplet size i
136 could suppress cloud precipitation process and extend cloud life time. The extension of the cloud life time might save the
137 potential that triggering the abnormal severe precipitation extreme events. This mechanism could partly explain the
138 precipitation degrading from light rain to severe extreme events (Fig. 6a) in the polluted EC region. Furthermore, the region
139 (west of 110 °E, south of 40 °N) of TP, a relative clean area in western China was selected as the reference area to
140 comparatively analyze the effects of aerosol pollution in hazy EC on regional precipitation change. we calculated
141 percentages of sites with negative frequency trends of light rain and positive trends of rainstorm events in total sites with
142 the positive and negative trends in haze over the EC and TP regions during the three interdecadal periods (1961-1980,
143 1971-2000, 1981-2010) (Fig. 6b). During past more than 5 decades, the light rain and rainstorm were receptively steady
144 declined and augmented in the polluted EC, while there were no obvious positive and negative variability trend of light rain
145 and rainstorm in the clean TP region (right panel of Fig. 6b).

146 **6. interannual anomalies between visibility and precipitation**

147 Figure 7 further verified the relation of interannual variability of regional visibility and precipitation in EC over recent years.
148 Regionally averaged, less light rain events and more rainstorms varied significantly from year to year in association with
149 enhanced aerosol levels with declining visibility over EC (Figs. 7a, 7b, 7c and 7d). , Taking summer months (June, July and
150 August) as examples, the 20-year running correlation coefficients of visibility and precipitation were presented in Figure 7e.
151 It is very interesting that the interannual variations of visibility and precipitation over EC were evolved from positive

152 correlations in the early 1960s to negative correlations in the 1970s and 1980s (Fig. 7e), reflecting the interannual variation
153 of the aerosol and precipitation interaction in changing climate.

154 In order to investigate the interannual variations of monthly correlation pattern between regional visibility with light rain and
155 extremely precipitation events in EC, we illustrated the cross-section of monthly anomalies of visibility and number of days
156 with light rain and rainstorm, in Figure 8.in. Through a comprehensive comparison of Figs. 8 a, b, c and d, we could find
157 significant positive correlation between visibility and light rain, indicating that the poor visibility suppressed light rain
158 frequency. Moreover, there was a significant difference between the changing extremely rainstorm and light precipitation
159 occurrences. The changes of large and extraordinary rainstorm frequency from 1960s to 1980s were not as prominent as
160 those at the latter period of 1990s, during which time visibility deteriorated remarkably, heavy and extremely heavy
161 rainstorm occurred frequently. Compared to other seasons, the influence effect of poor summer visibility was more
162 significant in EC, showing summertime disastrous rainfalls happened more often with less light rain over recent years. .

163 The increased atmospheric aerosol concentration may reduce the solar radiation to surface and decrease surface temperature.
164 forming a temperature inversion structure (Bollasina et al, 2011; Zhang et al., 2009; Bond et al., 2013; Bond et al., 2011;
165 Grant et al., 2014; Seinfeld et al., 2008). This temperature inversion structure with the stability of atmospheric boundary
166 layer provides an important condition for the frequent occurrence of haze events. The stable low-level structure also inhibits
167 the weak convection development of atmospheric boundary layer, reducing the formation of low-level clouds and weak
168 precipitation process. However, the strong dynamic convergence disturbance could destroy the stability of atmospheric
169 boundary layer and cause the formation and development of severe rainstorms.

170 To further clarify the relation between aerosols and light rain frequency, the light rain frequency distribution from 601
171 stations in July, 2013 is displayed (Fig.9). The light rain events have significantly declined in EC with high aerosol
172 concentrations have and but enhanced in the relative clean TP region (Fig. 9).

173 To reveal the relationship between aerosols and atmospheric vertical thermal structure, the correlation between surface $PM_{2.5}$
174 concentrations and atmospheric thermal structure in both polluted and clean areas in July, 2013 was investigated (Fig. 10).
175 The stations of Changsha and Hongjia located in Hunan and Zhejiang provinces in EC respectively were selected to
176 represent the less light rain region while those of Linzhi and Dingri of TP were selected to represent the high-frequency light
177 rain region. The correlation coefficient profiles between the observed surface daily $PM_{2.5}$ concentration and atmospheric
178 temperature profiles derived from high-resolution L-band sounding were calculated. The correlations at Changsha and
179 Hongjia stations (Figs.10a-b) show that the correlation between $PM_{2.5}$ and temperature profiles presented an "inverse phase"
180 pattern, reflecting the high aerosol concentrations in a thermal stable structure similar to temperature inversion layers with
181 "cold at low-layer and warm at upper-layer" in the EC. On the contrary, the correlations in Linzhi and Dingri stations in the
182 TP (Fig. 10c-d) indicate that an unstable atmospheric structure with "warm at low-layer and cold at upper-layer" with a

183 favorable condition for the occurrence and development of convection and light rain events in the TP.

184 **7. Physical connection between aerosols and precipitation**

185 According to the results of observation and modeling studies, the increased aerosol concentrations could reduce effective
186 particle radius and increased number concentration of cloud droplets (Khain et al., 2005; Van den Heever et al., 2006; Tao et
187 al., 2007; Altaratz et al., 2014). The increase of cloud droplets concentrations would delay raindrop formation, thereby
188 lessening light precipitation (Qian et al., 2009) for the negative correlation between aerosols and light precipitation in China
189 (Choi et al., 2008).

190 In order to further confirm the relationship between aerosols and cloud droplets, the cloud droplet data observed by aircraft
191 in north China during 2008-2010 were used. The vertical profiles of cloud droplets under different aerosol state obtained by
192 40 aircraft are shown in Figure 11. Aerosol Albrecht "cloud lifetime effect" was significant in the northern EC (Figs. 11a and
193 11b). Under the background of high aerosol concentrations, the cloud droplets sizes were smaller and increased slowly with
194 the increasing altitude (red profile of Fig. 11b). In addition, from the cloud base to 2000m, the cloud droplets size remained
195 less than 20 microns (Fig.11b); resulting in precipitation delay, in favor of cloud system development to form heavy rain
196 easily. Cloud droplets diameter enlarged quickly with the increase of height, and reached 30 microns easily to forming light
197 rain at 1000m altitude under low aerosol concentrations (green profile in Fig.11b). The aircraft observation analysis showed
198 that high aerosol concentrations could reduce cloud droplets size, increase cloud droplets concentrations, extend cloud
199 lifetime, which could restrict the light rain process.

200 **8. Discussion and conclusions**

201 Aerosols have complicated effects on clouds and precipitation, depending on many factors such as aerosol properties,
202 topography and meteorological conditions. The most previous investigations of aerosol impacts on clouds and precipitation
203 are primarily based on limited cases in relatively small spatial and temporal scales. The climate forcing of aerosols on
204 precipitation in large-scale region and physical causes remain uncertain. By using precipitation and visibility data of more
205 than 50 years, aircraft and surface aerosol data in recent years in China, the impacts of aerosol variations on interannual
206 variability of various precipitation intensities of precipitation events and their physical causes are investigated.

207 Accompanied with the frequent haze events in EC, the light rain frequency trend significantly decreased. Especially, since
208 the 1980s the extremely heavy precipitation event have occurred more frequently with an obvious transform from more light
209 rain to more frequent heavy rain and rainstorm. In the 1960s, the monthly visibility and light rain presented a significantly
210 positive correlation, while the visibility was in good condition. In recent 30 years, the dramatically increased aerosols
211 resulted in poor visibility, and the light rain frequency decreased obviously, and, heavy and extremely heavy rain occurred

212 more frequently.

213 The investigation of relation between aerosol concentrations and light rain frequency distributions in July, 2013 in China
214 shows that that the light rain appeared significantly low frequency in the EC region with high aerosol concentrations , and
215 but high-frequency in the relative clean region of Tibetan plateau presented significantly. High aerosol concentrations was
216 strongly correlated to warming low-level atmospheric to forming a stable structure suppressing the occurrence and
217 development of light rain events in EC. The aircraft measurements over the EC confirmed that the diameters of cloud
218 droplets decreased under high aerosol concentration condition, thereby inhibiting weak precipitation process.

219 The findings from this study have important implications for aerosol and precipitation interaction. The frequent haze events
220 in EC not only cause regional environment deterioration, but also induce the long-term change of regional water cycle with
221 the effect on regional climate change.

222

223 **Acknowledgements**

224 The study was supported by the National Key R & D Program Pilot Projects of China (JFYS2016ZY01002213;
225 2016YFC0203304), the National Natural Science Foundation of China (91544109; 91644223), and the project of
226 Environmental
227 Protection (HY14093355; 201509001) in the Public Interest and Chinese Third Tibetan Plateau Atmospheric Experiment
228 (GYHY201406001).

229 **Reference**

230 Altaratz, O., Koren, I., Remer, L. A. and Hirsch, E.: Review: Cloud invigoration by aerosols-coupling between microphysics
231 and dynamics, *Atmos. Res.*, 140-141, 38-60, 2014.

232 Bollasina, M. A., Ming, Y., and Ramaswamy, V.: Anthropogenic aerosols and the weakening of the South Asian summer
233 monsoon, *Science* 334, 502-505, 2011.

234 Bond, T. C., and Coauthors: Bounding the role of black carbon in the climate system: A scientific assessment, *J. Geophys.*
235 *Res.*, 118, 5380-5552, 2013.

236 Bond, T. C., Zarzycki, C., Flanner, M. G., and Koch, D. M.: Quantifying immediate radiative forcing by black carbon and
237 organic matter with the Specific Forcing Pulse, *Atmos. Chem. Phys.* 11, 1505-1525, 2011.

238 Choi, Y. S., Ho, C. H., Kim, J., Gong, D. Y. and Park, R. J.: The impact of aerosols on the summer rainfall frequency in
239 China, *J. Appl. Meteor Climatol.*47, 1802-1813, 2008.

240 Fan, J., Leung, L. R., Rosenfeld, D., Chen, Q., Li, Z., Zhang J. and Yan H.: Microphysical effects determine macrophysical
241 response for aerosol impacts on deep convective clouds. *Proc. Natl. Acad. Sci.*110, E4581-E4590, 2013.

242 Fan, J., Rosenfeld, D., Yang, Y., Zhao, C., Leung, L. R. and Li, Z.: Substantial contribution of anthropogenic air pollution to
243 catastrophic floods in southwest China, *Geophys. Res. Lett.* 42, 6066-6075, 2015.

244 Fu, Q., Johanson, C. M., Wallace, J. M. and Reichler, T.: Enhanced midlatitude tropospheric warming in satellite
245 measurements, *Science* 312, 1179, 2006.

246 Grant, L. D. and van den Heever, S. C.: Microphysical and dynamical characteristics of low-precipitation and classic
247 supercells, *J. Atmos. Sci.* 71, 2604–2624, 2014.

248 Guo X. L., Fu, D. H., Guo, X. and Zhang, C. M.: A case study of aerosol impacts on summer convective clouds and
249 precipitation over northern China, *Atmos. Res.* 142, 142-157, <http://dx.doi.org/10.1016/j.atmosres.2013.10.006>, 2014.

250 IPCC, In Summary for Policymakers, in *Climate Change 2007: The Physical Science Basis*, Contribution of Working Group
251 I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change. Cambridge: Cambridge University
252 Press, 1-18, 2013.

253 Khain, A. P., Rosenfeld, D. and Pokrovsky, A.: Aerosol impact on the dynamics and microphysics of deep convective clouds,
254 *Quart. J. Roy. Meteor. Soc.* 131, 2639-2663, 2005.

255 Koren, I., Martins, J. V., Remer, L. A. and Afargan, H.: Smoke invigoration versus inhibition of clouds over the Amazon,
256 *Science* 321, 946-949, 2008.

257 Lau, K. M. and Wu, H. T.: Detecting trends in tropical rainfall characteristics, 1979-2003, *Int. J. Climatol.*, 27, 979-988,
258 2007.

259 Levin, Z. and Cotton, W. R.: *Aerosol pollution impact on precipitation: A scientific review*, Springer, 386, 2009.

260 Li, Z., Niu, F., Fan, J., Liu, Y., Rosenfeld, D. and Ding, Y.: Long-term impacts of aerosols on the vertical development of
261 clouds and precipitation, *Nat. Geosci.* 4, 888–894, 2011.

262 Liu, S. C., Fu, C., Shiu, C. J., Chen, J. P. and Wu, F.: Temperature dependence of global precipitation extremes, *Geophys.*
263 *Res. Lett.* 36, L17702, doi: 10.1029/2009GL040218, 2009.

264 Qian, W., Fu, J., and Yan, Z.: Decrease of light rain events in summer associated with a warming environment in China
265 during 1961-2005, *Geophys. Res. Lett.* 34, L11705. DOI: 10.1029/2007GL029631, 2007.

266 Qian, Y., Gong, D. Y., Fan, J. W., Leung, L. R., Bennartz, R., Cheng, D. L. and Wang, W. G.: Heavy pollution suppresses
267 light rain in China: Observations and modeling, *J. Geophys. Res.* 114, D00K02, doi: 10.1029/2008JD011575, 2009.

268 Ramanathan, V., Chung, V. C. and Kim, D.: Atmospheric brown clouds: Impacts on South Asian climate and hydrological
269 cycle, *Proc. Natl. Acad. Sci.* 102, 5326–5333, 2005.

270 Ramanathan, V., Ramana, M. V. and Roberts, G.: Warming trends in Asia amplified by brown cloud solar absorption,
271 *Nature* 448, 575–578, 2007.

272 Rosenfeld, D. and Coauthors.: Flood or drought: How do aerosols affect precipitation? *Science* 321, 1309-1313, 2008.

273 Rosenfeld, D., Dai, J., Yu, X., Yao, Z., Xu, X., Yang, X. and Du, C.: Inverse relations between amounts of air pollution and
 274 orographic precipitation, *Science* 315,1396-1398, 2007.

275 Seinfeld, J.: Atmospheric science: Black carbon and brown clouds, *Nat Geosci* 1, 15–16, 2008.

276 Stevens, B. and Feingold, G.: Untangling aerosol effects on clouds and precipitation in a buffered system, *Nature* 461,
 277 607-613, 2009.

278 Tao, W. K., Li, X., Khain, A., Matsui, T., Lang, S. and Simpson, J.: Role of atmospheric aerosol concentration on deep
 279 convective precipitation: Cloud-resolving model simulations, *J. Geophys. Res.*, 112, D24S18, 2007.

280 Van den Heever, S. C., Carrió, G. G. Cotton, W. R. DeMott, P. J. and Prenni, A. J.: Impact of nucleating aerosol on Florida
 281 storms. Part I: Mesoscale simulations, *J. Atmos. Sci.*63, 1752-1775, 2006.

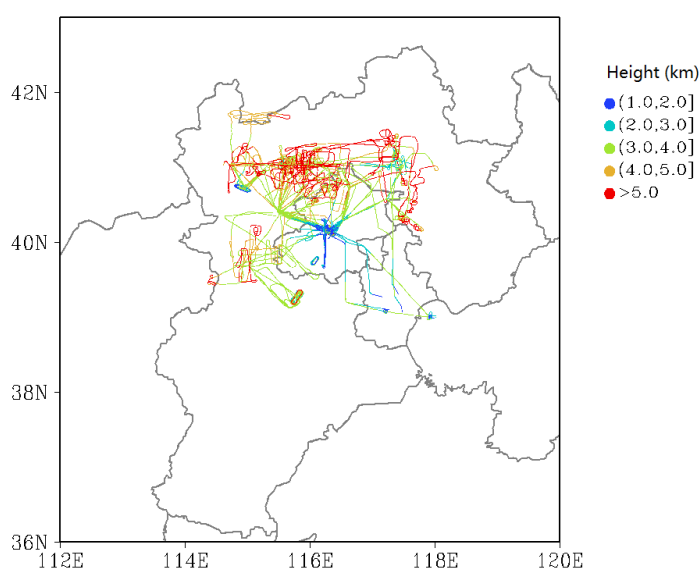
282 Wang, Y. and Zhou, L.: Observed trends in extreme precipitation events in China during 1961-2001 and the associated
 283 changes in large-scale circulation, *Geophys. Res. Lett.* 23, L09707, doi:10292005GL022574, 2005.

284 Zhai, P.M., Ren, F. M. and Zhang, Q.: Detection of trends in China's precipitation extremes, *Acta Meteorologica Sinica*, 57,
 285 208-215, 1999.

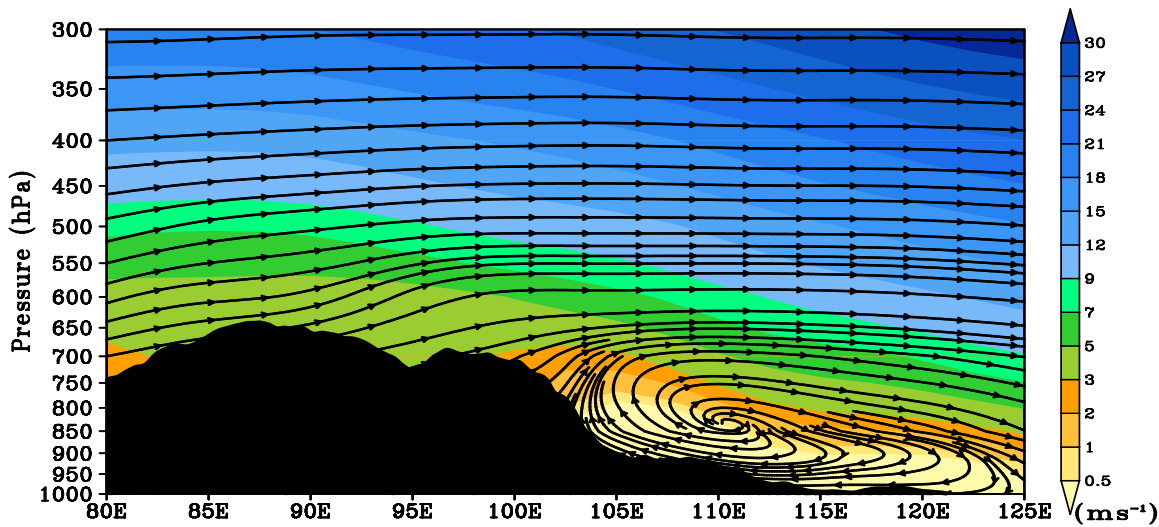
286 Zhang, H., Wang, Z. L., Guo, P. W., and Wang, Z. Z.: A modeling study of the effects of direct radiative forcing due to
 287 carbonaceous aerosol on the climate in East Asia, *Adv. Atmos. Sci.*, 26, 57–66, 2009.

288 Zhao, T. Liu, D., Zheng, X., Yang, L., Gu, X., Hu, J., Shu, Z., Chang, J., Wu, X.: Revealed variations of air quality in
 289 industrial development over a remote plateau of Southwest China: an application of atmospheric visibility data, *Meteorol*
 290 *Atmos Phys*, doi:10.1007/s00703-016-0492-7, 2016.

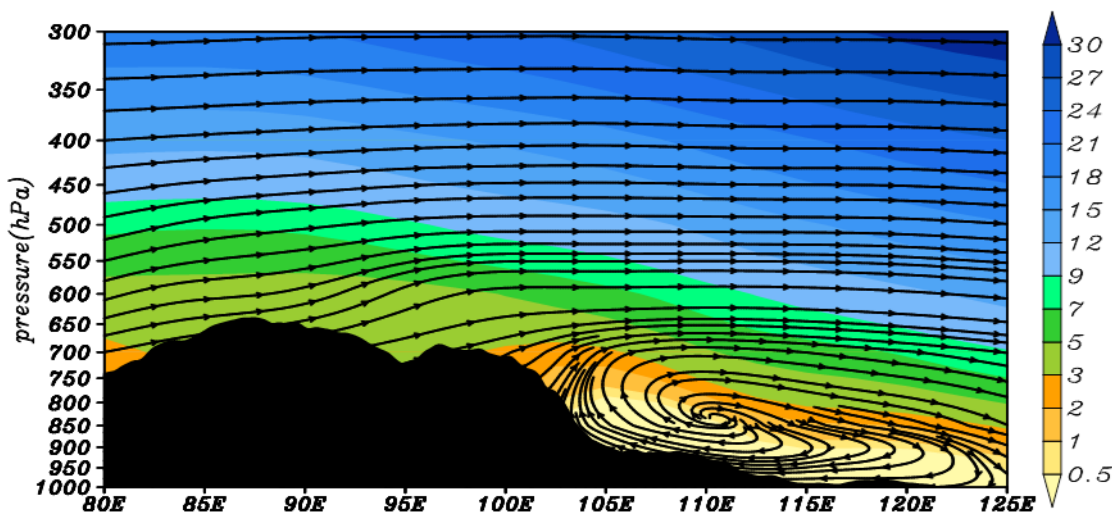
291
 292



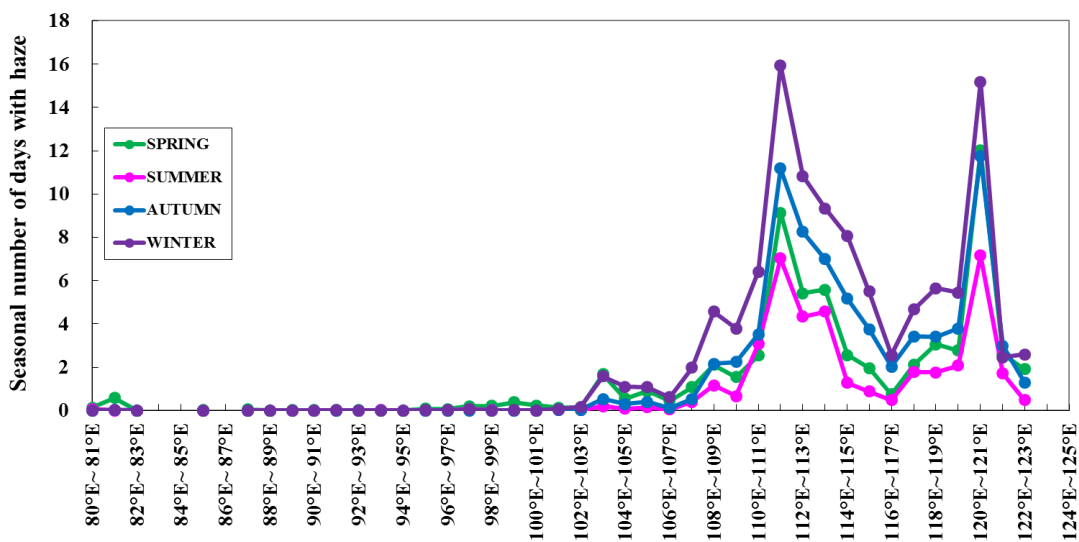
293
 294 **Figure 1 Area and tracks of 40 aircraft flights carried out in Beijing and its surrounding regions during aerosol-cloud experiment**
 295 **from 2008 to 2010 by the Beijing Weather Modification Office. China**



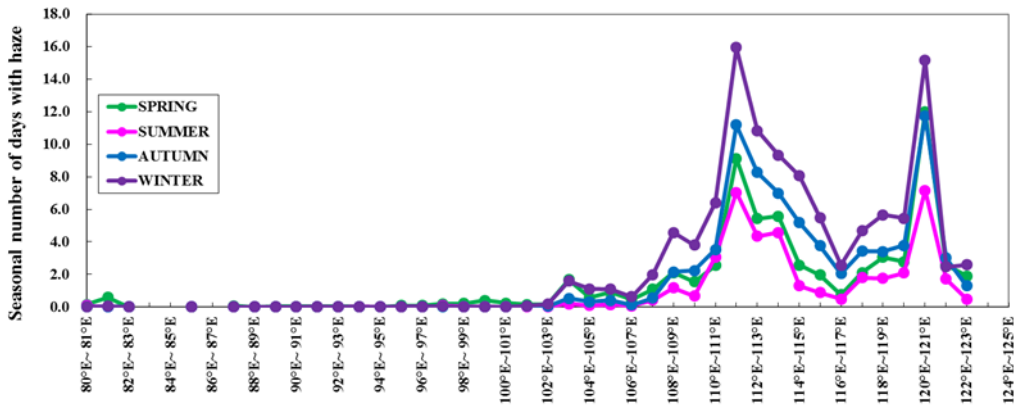
297



298



299



300

301

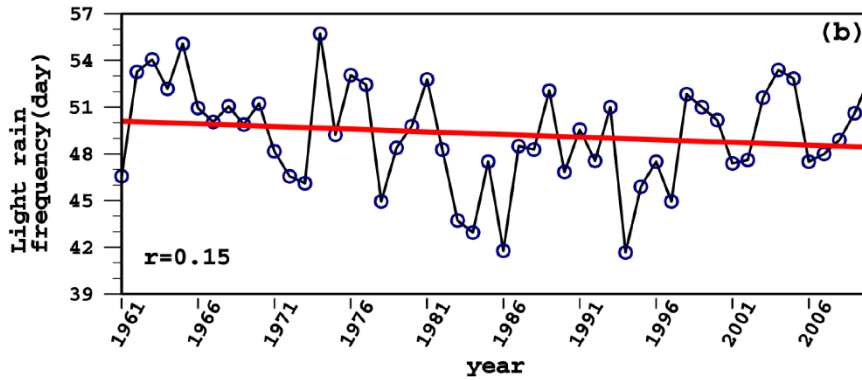
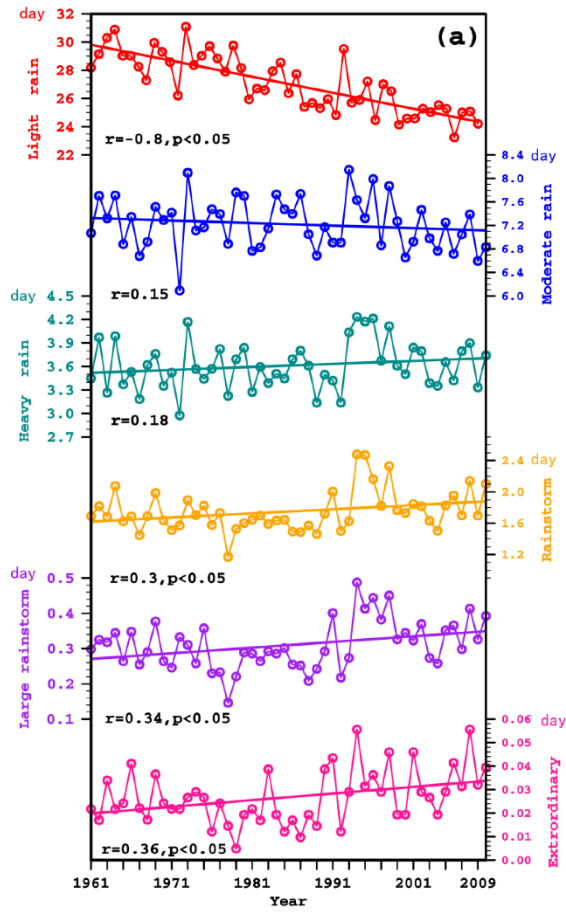
302

303

304

305

Figure 2. Cross sections of vertical circulations illustrated by stream lines (upper panel) with the horizontal wind speed (m s^{-1} ; color contours) and zonal variations of annual haze event frequency (lower panel) at 27°N - 41°N averaged in spring, summer, autumn and winter over 1961-2012. Note that near-surface vertical and horizontal winds are not illustrated well here due to north-south variations in the terrain and approximation of the location of the plateaus (black shaded area) in upper panel. All fields are for the annual-averages.



306

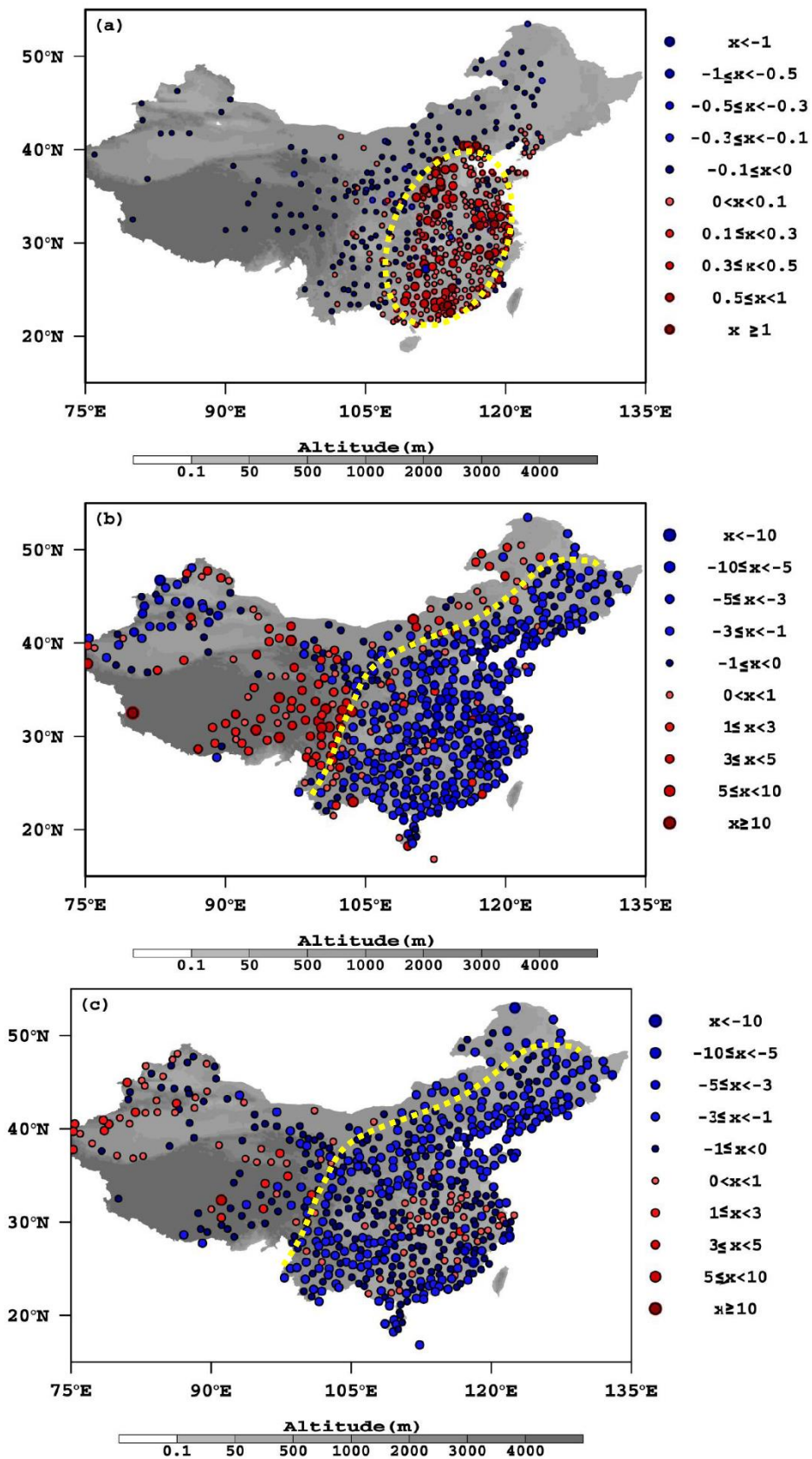
307 **Figure 3** Interannual variations with their anomalies (broken lines) and trends (straight lines) in (a) various precipitation
 308 intensities in the high aerosol concentration area in the EC region and (b) light rain in the relative clean area of Tibetan Plateau.

309

310

311

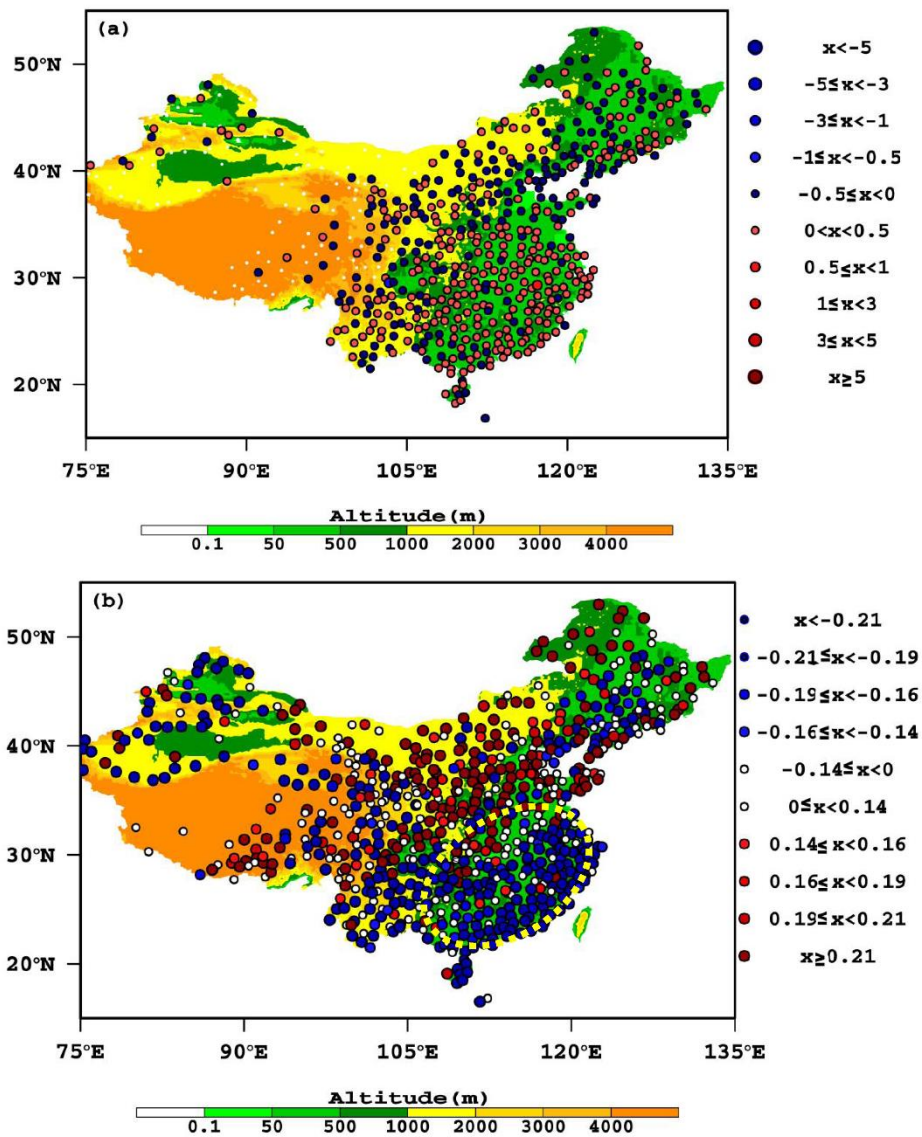
312



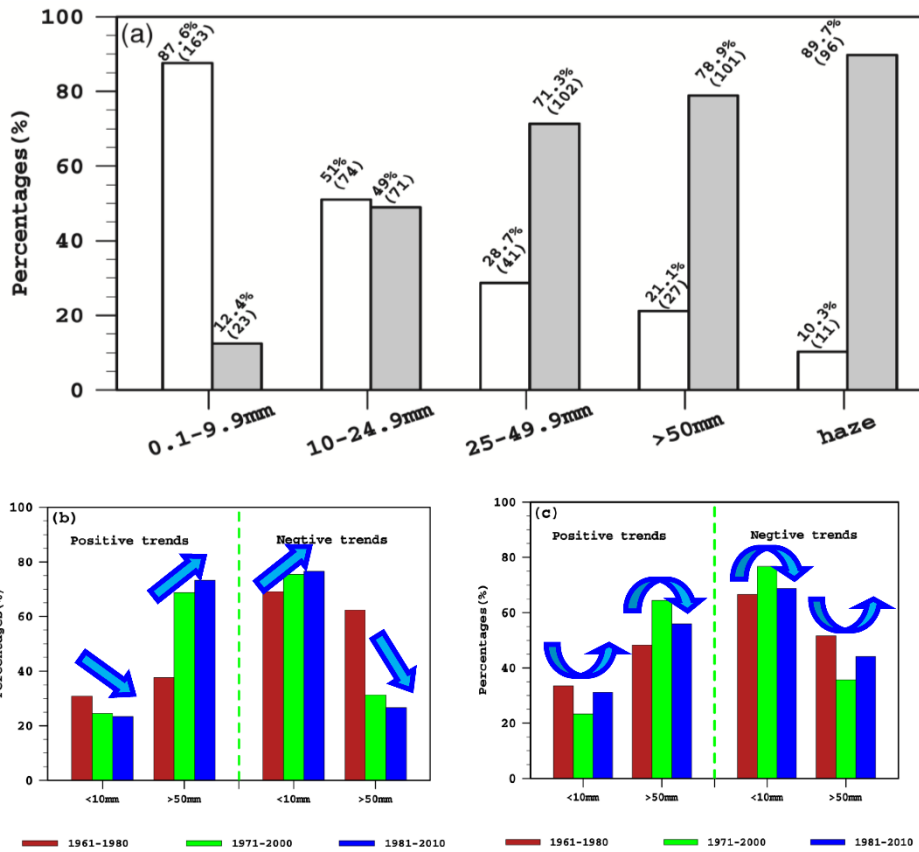
313

314 Figure 4 Distribution of interannual change trends (day per 10 years) in (a) haze frequency, (b) visibility and (c) light rain

315 frequency in summer in mainland China in 1961-2010. The yellow dash lines mark the borders of frequent haze area or the



319 Figure 5 The spatial distributions of (a) trends (day per 10 years) in summertime rainstorm frequency over 1961-2010 in China
 320 and (b) correlation coefficients between visibility and low cloud amount in summer of 1961-2010 With the yellow dash line
 321 marking the border of negative correlation area.

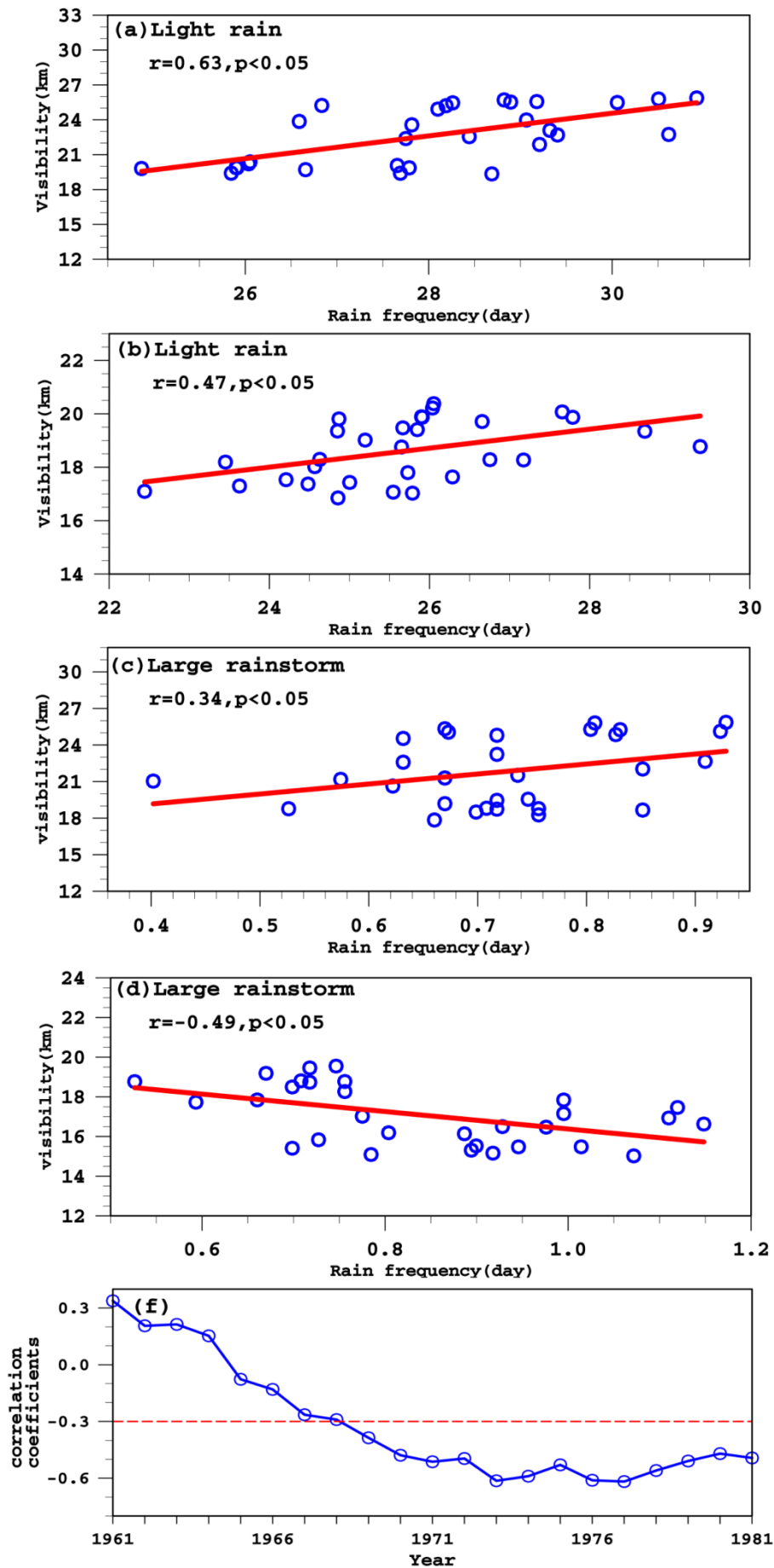


322

323 Figure 6 (a)percentages of stations with by positive and negative trends in number of day in various precipitation grades from
 324 1961 to 2010 over the EC stations ,(The percentages of sites with negative frequency trends of light rain and positive trends of
 325 rainstorm events in total sites with the and positive (left side) and negative (right side) trends in haze over (b) the EC and (c) TP
 326 regions during the three interdecadal periods (1961-1980, 1971-2000, 1981-2010). The arrows indicate the interdecadal change
 327 patterns.

328

329

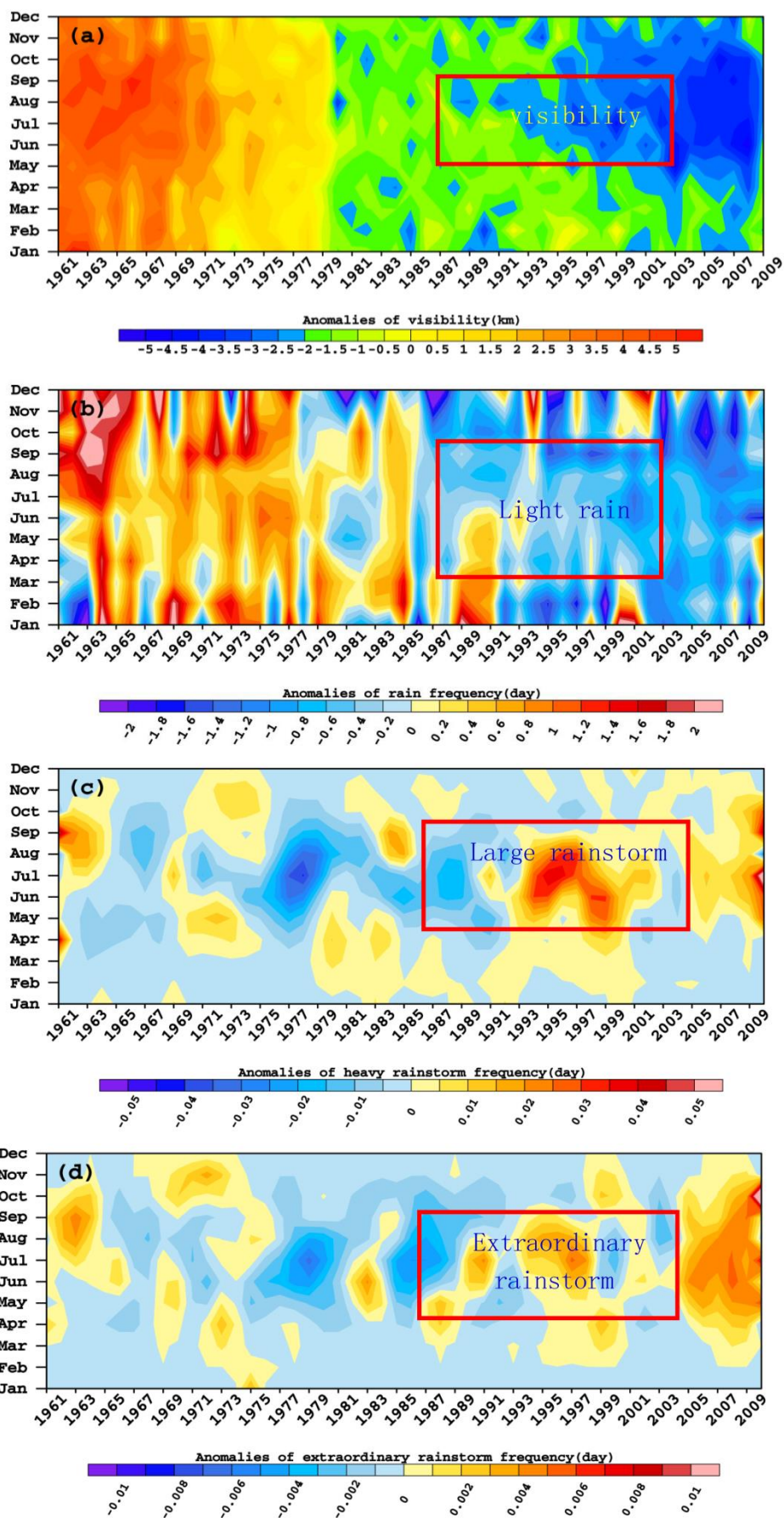


330

331 Figure 7 Correlation between summer average visibility (June, July and August) with light rain frequency (a) in 1961-1990, (b)

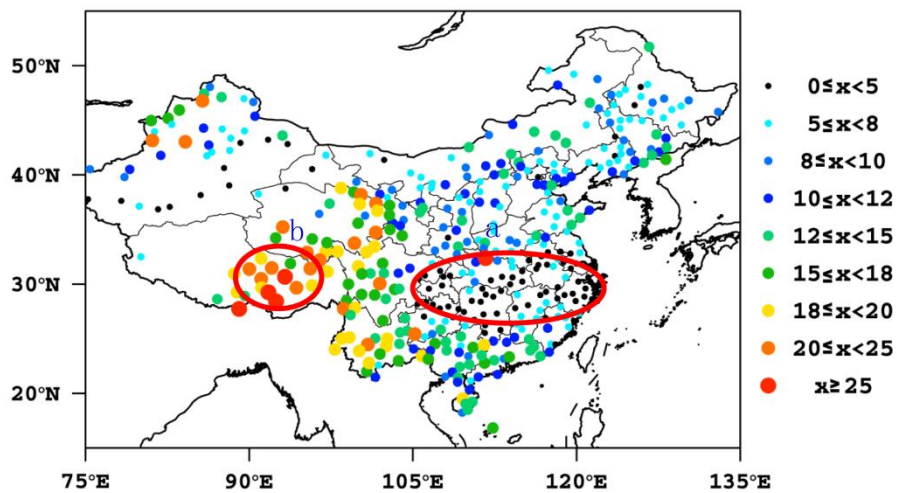
332 light rain frequency in 1981-2010, (c) extremely heavy rain event frequency in 1961-1990, and (d) extreme heavy rain event

333 frequency in 1981-2010; (e) The 20-year running correlation coefficients of visibility and precipitation over EC with dash line
 334 standing for the correlation passing the confidence level of 90%.



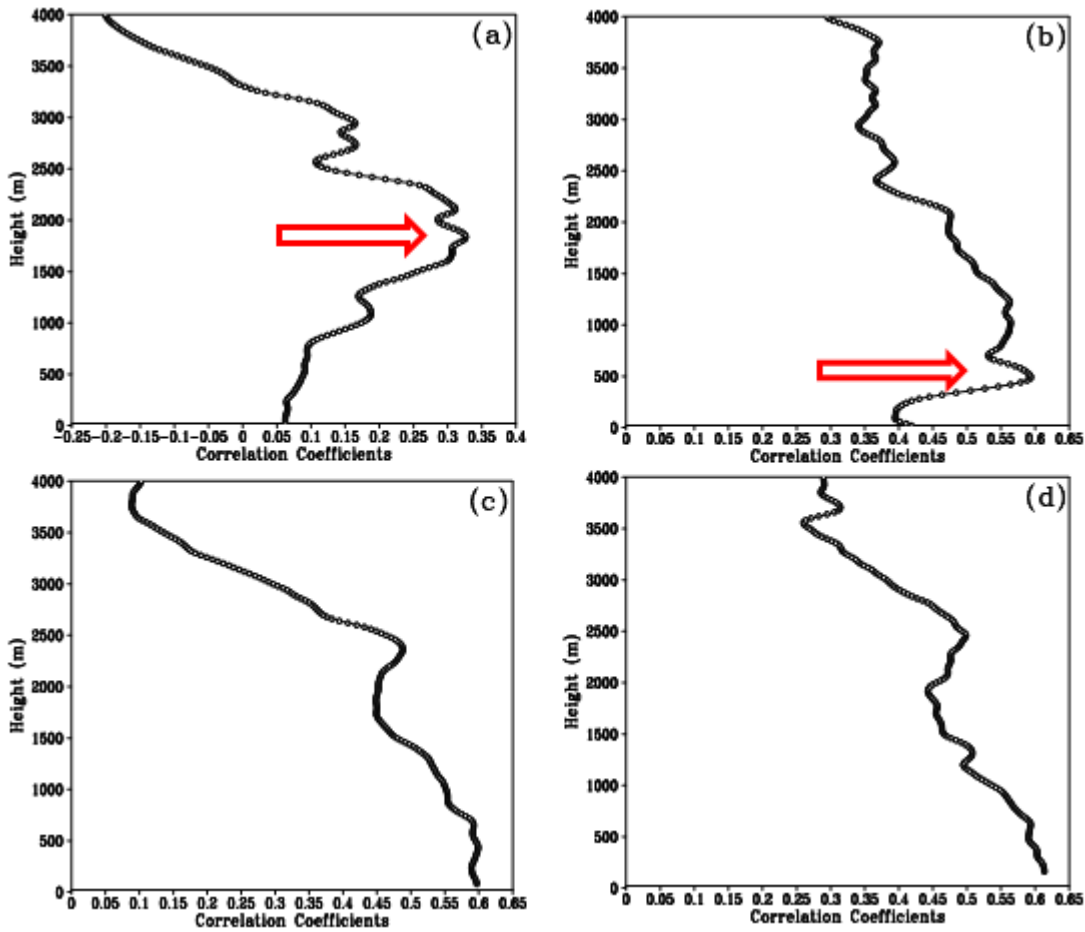
335
 336 Figure 8 Monthly anomalies of number of days with (a) visibility, (b) light rain, (c) heavy rain, and (d) extremely heavy rain

337 from 1961-2010 averaged over EC. The red rectangles mark the areas with significant changes.
338
339
340
341



342
343 **Figure 9** Light rain frequency distribution of 601 stations in China in July of 2013. The circled region on right is the low-frequency
344 light rain region in middle and downstream region in EC, and that on left is the high-frequency light rain region in relative clean
345 region over the TP.

346

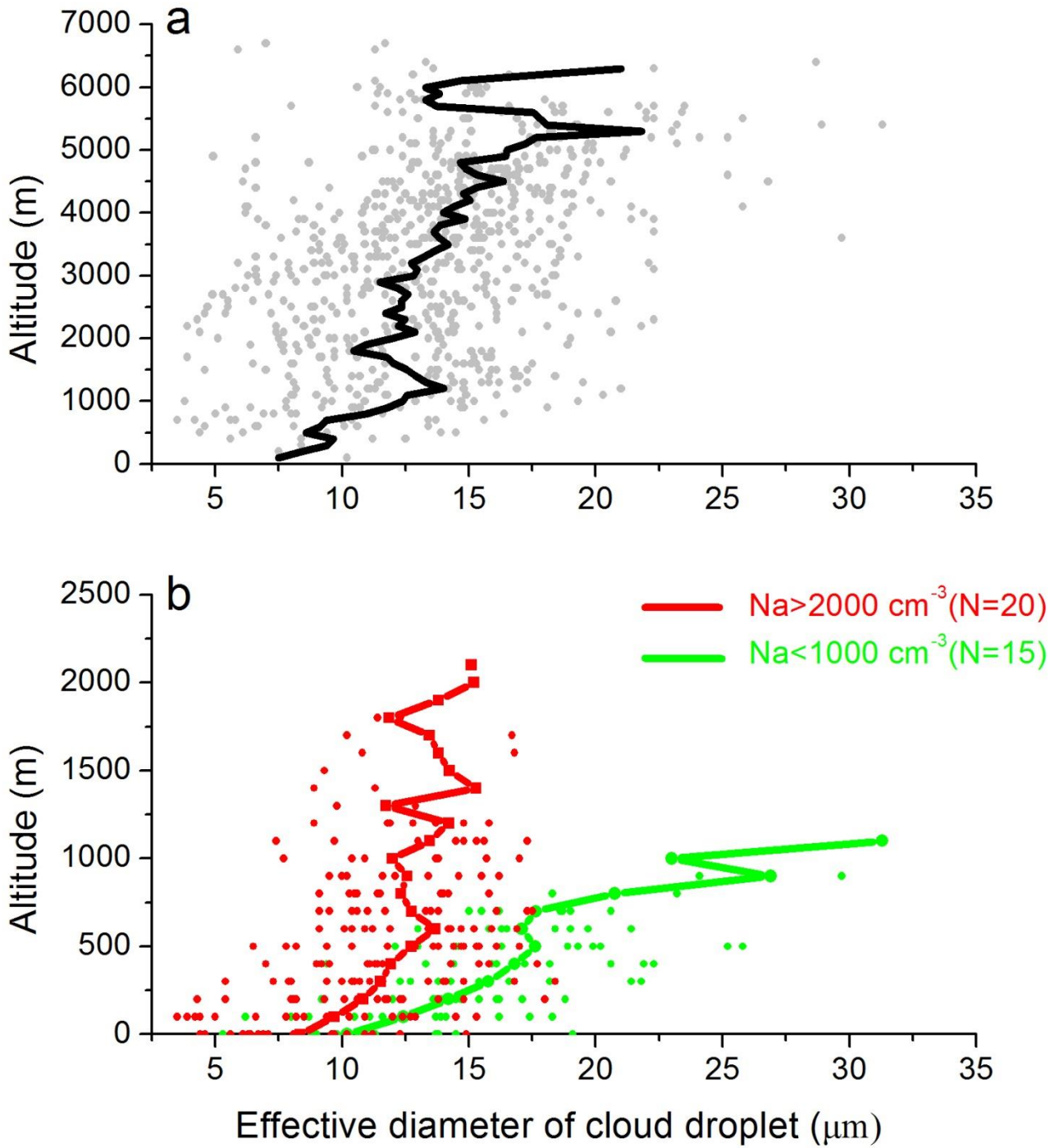


347

348 Figure 10 Correlation coefficient profiles between the surface $PM_{2.5}$ concentration (12 hour intervals) and atmospheric
 349 temperature from L-band sounding for representing low-frequency light rain regions at stations of (a) Changsha and (b)
 350 Hongjia in EC, and for representing high-frequency light rain regions at relative "clean area" at stations of (c) Linzhi and (d)
 351 Dingri over the TP with correlation coefficients of 0.21, 0.25 and 0.32 passing the confidence level of 90%, 95% and 99%.

352

353



354

355

356

Figure 11 the vertical profiles of (a) sampling cloud droplet size detected by 40 aircraft, and (b) changes and number and diameters of cloud droplets under low and high aerosol number concentrations.