



1 Source attribution of black carbon and its direct radiative forcing
2 in China

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17



18 **Abstract**

19 The source attributions for mass concentration, haze formation, transport, and
20 direct radiative forcing of black carbon (BC) in various regions of China are quantified
21 in this study using the Community Earth System Model (CESM) with a source-tagging
22 technique. Anthropogenic emissions are from the Community Emissions Data
23 System that is newly developed for the Coupled Model Intercomparison Project
24 Phase 6 (CMIP6). Over North China where the air quality is often poor, about 90% of
25 near-surface BC concentration is contributed by local emissions. 30% of BC
26 concentration over South China in winter can be attributed to emissions from North
27 China and 10% comes from sources outside China in spring. For other regions in
28 China, BC is largely contributed from non-local sources. We further investigated
29 potential factors that contribute to the poor air quality in China. During polluted days,
30 a net inflow of BC transported from non-local source regions associated with
31 anomalous winds plays an important role in increasing local BC concentrations.
32 BC-containing particles emitted from East Asia can also be transported across the
33 Pacific. Our model results show that emissions from inside and outside China are
34 equally important for the BC outflow from East Asia, while emissions from China
35 account for 7% of BC concentration and 25% in column burden in western United
36 States in spring. Radiative forcing estimated shows that 66% of the annual mean BC
37 direct radiative forcing (2.3 W m^{-2}) in China results from local emissions, and the
38 remaining 34% are contributed by emissions outside of China. Efficiency analysis
39 shows that reduction in BC emissions over eastern China could benefit more on the
40 regional air quality in China, especially in winter haze season.



41 1. Introduction

42 Black carbon (BC), as a component of atmospheric fine particulate matter
43 (PM_{2.5}), is harmful to human health (Anenberg et al., 2011; Janssen et al., 2012). In
44 addition to its impact on air quality, as the most efficient light-absorbing
45 anthropogenic aerosols, BC is thought to exert a substantial influence on climate
46 (Bond et al., 2013; IPCC, 2013; Liao et al., 2015). It can heat the atmosphere through
47 absorbing solar radiation (Ramanathan and Carmichael, 2008), influence cloud
48 microphysical and dynamical processes (Jacobson, 2006; McFarquhar and Wang,
49 2006), and reduce surface albedo through deposition on snow and ice (Flanner et al.,
50 2007; Qian et al., 2015).

51 Due to accelerated urbanization and rapid economic growth, emissions of BC in
52 China increased dramatically during recent decades. It contributed to about one
53 fourth of the global emissions of BC in recent decades (Bond et al., 2007). Strong
54 emissions lead to high concentrations of BC over China. Zhang et al. (2008) collected
55 aerosol samples at 18 stations spread over China during 2006 and reported BC
56 concentrations in a range of 9–14 $\mu\text{g m}^{-3}$ at urban sites, 2–5 $\mu\text{g m}^{-3}$ at rural sites, and
57 about 0.35 $\mu\text{g m}^{-3}$ at remote background sites. BC also exerts significant positive
58 direct radiative forcing (DRF) at the top of the atmosphere (TOA) in China. Using the
59 Regional Climate Chemistry Modeling System (RegCCMs), Zhuang et al. (2013)
60 reported an annual mean BC DRF of 2–5 W m^{-2} at TOA over eastern China and
61 about 6 W m^{-2} over Sichuan Basin in year 2006. Li et al. (2016) also showed a strong
62 DRF of BC over the North China Plain and Sichuan Basin in most seasons except for
63 spring when the strongest BC DRF with values of 4–6 W m^{-2} shifted to southern
64 China.

65 BC is the product of incomplete combustion of fossil fuels, biofuels, and open
66 burning, such as forest and grassland fires and agricultural waste burning on fields. In
67 the atmosphere the average lifetime of BC is only a few days, due to both wet
68 removal and dry deposition, which is much shorter than that of long-lived greenhouse
69 gases. In addition, BC lifetime is region dependent. BC in East Asia has a shorter
70 lifetime than the global mean value due to a faster regional removal (H. Wang et al.,



71 2014). BC emission reductions may benefit both mitigation of global climate change
72 and regional air quality (Shindell et al., 2012; Bond et al., 2013; Smith and Mizrahi,
73 2013), especially in East Asia where fuel combustion emits substantial BC along with
74 other pollutant species. Many previous observational and/or modeling studies have
75 examined the source sector contributions of BC over China (Zhuang et al., 2014;
76 Y.-L. Zhang et al., 2015; Li et al., 2016). They found that residential heating and
77 industry sectors were the largest contributors to BC concentrations in China, while
78 biomass burning emissions from outside China were important to BC in western
79 China. An effective BC reduction in a receptor region would require knowing not only
80 the source sector that contributes the most to BC levels, but also the source
81 contributions from various locations within and outside the region. However, very few
82 previous studies have focused on the source attribution of BC concentrations in
83 various regions of China. Li et al. (2016) examined the contributions of emissions
84 inside and outside China to BC over China (with only two source regions) but did not
85 divide the source contributions from different regions inside China.

86 Pollution levels also show substantial daily to weekly variation. In recent years,
87 extreme wintertime hazy conditions occurred frequently in China and caused serious
88 air pollution, affecting more than half of the 1.3 billion people (Ding and Liu, 2014).
89 During one winter haze episode in 2013, BC concentrations increased up to about 20
90 and $8 \mu\text{g m}^{-3}$ in Xi'an and Beijing over northern China, and 6 and $4 \mu\text{g m}^{-3}$ in
91 Guangzhou and Shanghai over southern China, respectively (Y.-L. Zhang et al.,
92 2015). The transport of pollutants from upwind was reported to be one of the most
93 important contributors to local high aerosol concentrations during haze days (L. T.
94 Wang et al., 2014; Y. Yang et al., 2016). L. T. Wang et al. (2014) found that emissions
95 from northern Hebei and Beijing-Tianjin were the major contributor to particulate
96 matter ($\text{PM}_{2.5}$) pollution in Shijiazhuang in January 2013. Yang et al. (2016) confirmed
97 a connection between wind fields and $\text{PM}_{2.5}$ concentrations during winter hazy days
98 through model simulations and statistical analysis. They also found that weakened
99 winds contributed to increases in winter aerosol concentrations and hazy days over
100 eastern China during recent decades. As a chemically inert species, atmospheric BC



101 is a good tracer to investigate the source region contributions from local and non-local
102 emissions during polluted conditions that are related to long-range transport.

103 BC particles originating from East Asia can also be transported across the North
104 Pacific, reaching North America (Hadley et al., 2007; Ma et al., 2013a; Matsui et al.,
105 2013; H. Wang et al., 2014; Yang et al., 2015). Matsui et al. (2013) simulated outflow
106 of BC from East Asia using the Community Multiscale Air Quality (CMAQ) model and
107 found that anthropogenic emissions from China, biomass burning emissions from
108 Southeast Asia, and biomass burning emissions from Siberia and Kazakhstan
109 contributed 61%, 17%, and 6%, respectively, to the eastward BC flux at 150°E
110 averaged over 2008–2010. Hadley et al. (2007) estimated the trans-Pacific transport
111 of BC during April of 2004 using the Chemical Weather Forecast System (CFORS)
112 model and reported that, across 130°W, 75% of BC transported into North America
113 originated from Asia. Huang et al. (2012) simulated BC using the Sulfur Transport
114 and Deposition Model (STEM), and found emissions outside North America
115 contributed to 30–80% of column BC over North America in summer 2008. H. Wang
116 et al. (2014) examined the long-term (1995–2005) average global source-receptor
117 relationship of BC and found that BC emitted from the entire East Asia only contribute
118 less than 5% to the total BC burden in North America, although the contribution is up
119 to 40% near the west coast region. Few studies have examined the outflow from East
120 Asia and inflow into North America contributed from source regions in and outside
121 China. In addition, the emissions of BC from China increased dramatically during the
122 last few years, with the annual total anthropogenic emissions estimated to have
123 almost doubled in year 2014 compared to year 2000, shown in the newly developed
124 Community Emissions Data System (CEDS; Hoesly et al. 2016). Therefore, the
125 long-range transport of BC and source-receptor relationships could be quite different
126 from previous studies.

127 Due to its warming effect in the climate system, BC is potentially important for
128 climate mitigation and has drawn much attention recently. Source attribution of the
129 direct radiative effect of BC is likely to be different from that of near-surface
130 concentration and column burden due to the dependence of radiative forcing on the



131 vertical distribution of BC and its mixing state with other species that are influenced
132 by different regional sources. In this study, we use the Community Earth System
133 Model (CESM) with improved representations of aerosol transport and wet removal
134 (H. Wang et al., 2013) and a BC source-tagging technique (H. Wang et al., 2014).
135 Anthropogenic emissions from the newly developed CEDS inventory (Hoesly et al.,
136 2016), as released for the Coupled Model Intercomparison Project Phase 6 (CMIP6),
137 are used to examine the source attributions for mass concentration, long-range
138 transport, and direct radiative forcing of BC in various regions of China. We aim to
139 quantify: (1) source region contributions to concentrations of BC over various
140 receptor regions in China; (2) contributions to changes in BC concentrations under
141 polluted conditions; (3) source contributions to trans-boundary and trans-Pacific
142 transport of BC; and (4) source contributions to direct radiative forcing of BC in China.

143 The CESM model, emissions, and numerical experiment are described in
144 Section 2. Section 3 provides evaluation of the simulated concentration and aerosol
145 absorption optical depth of BC in China. Section 4 investigates source contributions
146 to near-surface concentrations, long-range transport and direct radiative forcing of BC
147 over various receptor regions using the BC source-tagging technique in CESM.
148 Section 5 summarizes these results.

149

150 **2. Methods**

151 We simulate the evolution and direct radiative forcing (DRF) of BC using CESM
152 version 1.2 (Hurrell et al., 2013). The atmospheric model in CESM is version 5 of the
153 Community Atmosphere Model (CAM5), with horizontal grid spacing of 1.9° latitude
154 by 2.5° longitude and 30 vertical layers ranging from the surface to 3.6 hPa used in
155 this study. The model treats the properties and processes of major aerosol species
156 (sea salt, mineral dust, sulfate, black carbon, primary organic matter and secondary
157 organic aerosol) using a three-mode modal aerosol module (MAM3), in which aerosol
158 size distributions are represented by three lognormal modes: Aitken, accumulation,
159 and coarse modes. BC is emitted to the accumulation mode. Mass mixing ratios of
160 different aerosol species and the number mixing ratio are predicted for each mode. A



161 more detailed description of the MAM3 representation can be found in Liu et al.
162 (2012). The improved representations of processes related to convective transport
163 and wet scavenging of aerosols (H. Wang et al., 2013) are also included in this newer
164 version of the CESM/CAM5. Aerosol optical properties are parameterized in the
165 model according to Ghan and Zaveri (2007). Besides the default total radiative fluxes
166 calculated online, the CESM model can also calculate a separate set of radiative
167 fluxes without considering the contribution from one or more specific components.
168 The direct radiative forcing of those excluded species could then be calculated as the
169 differences of these two sets of radiative fluxes (Ghan, 2013).

170 Anthropogenic emissions used in this study are from the CEDS dataset, as
171 released for the CMIP6 model experiments (Hoesly et al. 2016). This newly released
172 emission inventory includes aerosol (black carbon, organic carbon) and aerosol
173 precursor and reactive compounds (sulfur dioxide, nitrogen oxides, ammonia, carbon
174 monoxide, and non-methane volatile organic compounds). The emissions are
175 provided at monthly resolution for each year of 1750–2014 on a $0.5^\circ \times 0.5^\circ$ grid and
176 include agricultural, energy, industry, residential, international shipping, solvents,
177 surface transportation, waste treatment, and aircraft sectors. The biomass burning
178 emissions used in this study are also developed for CMIP6 based on Global Fire
179 Emission Database (GFED) version 4, Fire Model Intercomparison Project (FireMIP),
180 visibility-observations and Global Charcoal Database (GCD) data (van Marle et al.
181 2016).

182 Figure 1a shows the horizontal spatial distribution of annual emissions of BC
183 averaged over the most recent 5 years (2010–2014) and the seven geographical
184 source regions tagged in continental China, including North China (NC), South China
185 (SC), Southwest China (SW), Central-West China (CW), Northeast China (NE),
186 Northwest China (NW), and Tibetan Plateau (TP). Figure 1b summarizes the total
187 seasonal BC emissions in each of these source regions. North China has the largest
188 annual emissions of BC in China, with maximum emission larger than 1.2 g C m^{-2}
189 year^{-1} and a regional total emission of $1089 \text{ Gg C year}^{-1}$ (44% of total emissions from
190 continental China). Annual emissions of BC also have large values over South China



191 and Southwest China, with maximum values in the range of 0.8–1.2 g C m⁻² year⁻¹,
192 followed by Central-West China and Northeast China. Over the less economically
193 developed Northwest China and remote region Tibetan Plateau, emissions of BC are
194 much lower than other regions in China. The seasonal mean emissions of BC also
195 show the same spatial pattern as the annual means. BC had the largest emissions
196 over North China, South China, and Southwest China in all seasons, among which
197 emissions are strongest in December-January-February (DJF), especially over North
198 China, resulting from domestic heating. The total seasonal emissions of BC in
199 continental China are 797, 586, 537, and 577 Gg C in DJF, March-April-May (MAM),
200 June-July-August (JJA), and September-October-November (SON), respectively,
201 which add up to a total annual BC emissions of 2497 Gg C averaged over years
202 2010–2014. The DJF emissions account for 26–35% of annual total whereas
203 emissions in JJA only account for 17–24% over the seven source regions in
204 continental China.

205 An explicit BC source tagging capability was originally implemented in CAM5 by
206 H. Wang et al. (2014), through which emissions of BC from independent source
207 regions and/or sectors can be explicitly tracked. This method quantifies the source–
208 receptor relationships of BC in any receptor region within a single model simulation
209 without perturbing emissions from individual source regions or sectors. R. Zhang et
210 al. (2015a,b) used this method to quantify the source attributions of BC in western
211 North America, Himalayas, and Tibetan Plateau. The same BC source tagging
212 technique is implemented to a newer model version (CAM5.3) and applied in this
213 study to quantify the source attributions of concentration, transport and direct
214 radiative forcing of BC in various regions of China. BC emissions (anthropogenic plus
215 biomass burning) from seven geographical source regions, including North China,
216 South China, Southwest China, Central-West China, Northeast China, Northwest
217 China, Tibetan Plateau in China, and from rest of the world (RW) are tagged.
218 Transport and physics tendencies are calculated separately for each tagged BC in
219 the same way as the original BC simulation in CESM. We choose the seven individual
220 regions (North China, South China, Southwest China, Central-West China, Northeast



221 China, Northwest China, and Tibetan Plateau) and all seven regions combined
222 (hereafter continental China) as receptor regions in this study to examine the
223 source-receptor relationships of BC. While all emissions, including sulfur dioxides,
224 organic carbon and BC, were used in the model simulation, tagging was only applied
225 to BC emissions.

226 The CAM5 simulation is performed at $1.9^\circ \times 2.5^\circ$ horizontal grid spacing using the
227 specified-dynamics mode (Ma et al., 2013b), in which large-scale circulations (i.e.,
228 horizontal winds) are nudged to 6-hourly reanalysis data from the Modern Era
229 Retrospective-Analysis for Research and Applications (MERRA) reanalysis data set
230 (Rienecker et al., 2011) with a relaxation time scale of 6 hours (K. Zhang et al., 2014).
231 The use of nudged winds allows for a more accurate simulation so that the key role of
232 large-scale circulation patterns matches observations over the specified years. The
233 simulation is run from year 2009 to 2014, with both time-varying aerosol emissions
234 and meteorological fields. The first year is for spin-up and the last five years are used
235 for analysis.

236

237 **3. Model evaluation**

238 The simulations of aerosols, especially BC, using CAM5 have been evaluated
239 against observations including aerosol mass and number concentrations, aerosol
240 optical properties, aerosol deposition, and cloud-nucleating properties in many
241 previous studies (e.g., Liu et al., 2012; H. Wang et al., 2013; Ma et al., 2013b; Jiao et
242 al., 2014; Qian et al., 2014; R. Zhang et al., 2015a,b). Here we focus on the
243 evaluation of model performance in China using measurements of near-surface BC
244 concentrations, aerosol index derived from satellite, and aerosol absorption optical
245 depth from the Aerosol Robotic Network (AERONET).

246 **3.1 Near-surface mass concentrations and column burden of BC**

247 Figure 2 presents spatial distributions of simulated seasonal mean near-surface
248 concentrations and column burden of BC, both of which show a similar spatial pattern
249 to emissions of BC (Figure 1a) with the largest values over North China and the
250 lowest values over Northwest China and Tibetan Plateau. Near-surface model results



251 are taken to be the lowest model layer (from surface to 993 hPa in average). Among
252 all seasons, DJF has the highest BC levels, with values in the range of 6–12, 2–8,
253 and 1–8 $\mu\text{g m}^{-3}$ for near-surface concentrations and 5–9, 3–7, 2–9 mg m^{-2} for column
254 burden over North, South, and Southwest China, respectively. In contrast, JJA has
255 the lowest BC concentrations over China due to the lower emissions and larger wet
256 scavenging associated with East Asian summer monsoon (Lou et al., 2016).
257 Averaged over continental China, near-surface BC concentrations are 2.2, 1.1, 0.8,
258 1.3 $\mu\text{g m}^{-3}$ in DJF, MAM, JJA, and SON, respectively, with seasonal variability of 38%.
259 The column burden of BC shows smaller seasonal variability (26%), with
260 area-weighted average of 1.9, 1.4, 1.1, and 1.3 mg m^{-2} in DJF, MAM, JJA, and SON,
261 respectively, in China. It suggests that, besides domestic emissions in China, there
262 are other BC sources from outside China contributing significantly to BC
263 concentrations in the column. The magnitude, spatial distribution, and seasonal
264 variations of simulated near-surface BC concentrations over China are similar to
265 those in Fu et al. (2012) and X. Wang et al. (2013) using Intercontinental Chemical
266 Transport Experiment-Phase B (INTEX-B) emission inventory (Zhang et al., 2009)
267 and those in Li et al. (2016) using Hemispheric Transport of Air Pollution (HTAP)
268 emission inventory (Janssens-Maenhout et al., 2015) together with a global chemical
269 transport model.

270 The simulated near-surface BC concentrations are evaluated here using
271 measurements at 14 sites of the China Meteorological Administration Atmosphere
272 Watch Network (CAWNET) (Zhang et al., 2012). The locations of CAWNET sites are
273 shown in Figure S1a. The observational data include monthly BC concentrations in
274 years 2006–2007. Note that the simulated BC concentrations are for years 2010–
275 2014. Figure 3a compares the simulated seasonal mean near-surface BC
276 concentrations with those from CAWNET observations and Table S1 summarizes the
277 comparison in different regions, using modeled values from the grid cell containing
278 each observation site. Simulated BC concentrations at most sites are within the range
279 of one third to 3 times of observed values, except for Dunhuang (94.68°E, 40.15°N)
280 and Lhasa (91.13°E, 29.67°N) sites over western China, where BC concentrations



281 appear to be underestimated in the model (up to 20 times lower). Over North China,
282 simulated concentrations are similar to observations in DJF, but underestimated in
283 other seasons. Over South China, the simulations do not have large biases compared
284 to the observed BC. However, simulated BC is underestimated in all seasons over
285 Southwest, Central-West, Northeast, Northwest China, and Tibetan Plateau.
286 Compared to the CAWNET data, the modeled near-surface BC concentrations have
287 a normalized mean bias (NMB) of -53% . Note that anthropogenic BC emissions went
288 up by a factor of 1.18 between 2006–2007 and 2010–2014. An emissions adjusted
289 comparison would result in an even larger underestimation. There are several
290 reasons that might cause low bias in this comparison. Liu et al. (2012) and H. Wang
291 et al. (2013) have previously found underestimation of BC concentrations over China
292 in CAM5 model and suggested the BC emissions may be significantly
293 underestimated. Using the global chemical transport model GEOS-Chem together
294 with emissions in 2006, Fu et al. (2012) found the simulated BC concentrations in
295 China were underestimated by 56%. With HTAP emissions at the year 2010 level, Li
296 et al. (2016) showed a low bias of 37% in simulated BC concentration in China. Large
297 wet removal rate and short lifetime of aerosols, resulting from the too frequent
298 liquid-containing cloud and fast precipitation scavenging in CAM5, also lead to the
299 lower concentrations of BC (Wang et al., 2011; Liu et al., 2012; H. Wang et al., 2013).

300 Another potential cause for a bias in this comparison is spatial sampling bias.
301 Half of the CAWNET sites are located in urban areas, which will tend to have high
302 values near sources, whereas the modeled values represent averages over large grid
303 cells (R. Wang et al., 2014), as further discussed below.

304 The model captures well the spatial distribution and seasonal variation of BC
305 concentrations in China, having a statistically significant correlation coefficient of
306 $+0.58$ between simulated and observed seasonal BC concentrations between
307 modeled and CAWNET values.

308 **3.2 Aerosol absorption optical depth of BC**

309 To evaluate the simulated aerosol absorption optical depth (AAOD) of BC, the
310 AAOD data from AERONET (Holben et al., 2001) are used here. The locations of



311 AERONET sites in China are shown in Figure S1b. The observed AAOD are
312 averaged over years of 2010–2014 over 7 sites and 2005–2010 over 3 sites with data
313 available. Most AERONET sites are over eastern and central China. AAOD of BC at
314 550nm are calculated by interpolating AAOD at 440 and 675 nm and removing AAOD
315 of dust from the retrieved AERONET AAOD following Bond et al. (2013). Figure 3b
316 compares the observed and simulated seasonal mean AAOD of BC at 550nm and
317 Table S2 summarizes the comparisons in different regions. The model has a low bias
318 in simulating AAOD of BC in China, smaller than the bias in near-surface
319 concentrations, with a NMB of –16% (Figure 3a). As is the case with surface
320 concentrations, this bias could be due to model issues, such as BC transport or
321 optical parameterization; an underestimate in emissions; or spatial sampling bias,
322 although spatial sampling bias is likely to be less important for the BC column than for
323 surface concentrations. Simulated AAOD of BC are within the range of one third to 3
324 times of observed values at most sites, with the spatial distribution and seasonal
325 variation broadly captured by the model. Note that all but one of the observations are
326 located in the North and South China regions, and simulated BC AAOD are, on
327 average, similar to observations there. The AAOD from one observation site in
328 Central-West China is higher than the modeled value.

329 Figure 4 shows the spatial distribution of simulated seasonal mean AAOD of total
330 aerosols and Aerosol Index (AI) derived from Ozone Monitoring Instrument (OMI)
331 measurements over years of 2010–2014. AI is a measure of absorbing aerosols
332 including BC and dust. Compared to satellite AI data, the model roughly reproduces
333 spatial distribution of total AAOD in China, with large values over North, South, and
334 Southwest China in all seasons. AI from derived from Total Ozone Mapping
335 Spectrometer (TOMS) measurements also show similar spatial pattern as simulated
336 AAOD (Figure S2). It should be noted that, besides BC, dust particles also largely
337 contribute to AI and produces large AI values over Northwest China.

338 To examine the potential model bias more broadly we compared the difference of
339 AAOD and AI western China and eastern China (Fig. 4). Averaging AI and AAOD
340 broadly over eastern and western China, we find that AAOD/AI is 0.048 over eastern



341 China and 0.031 over western China. If we assume the simulated AAOD do not have
342 large bias over eastern China compared to observations, then this difference hints at
343 a possible underestimation of BC column burden in the model over the western
344 regions. It is somewhat difficult to draw a firm conclusion, however, given the likely
345 differential role of dust, and model biases modeling dust, and possible biases in
346 satellite derived AI values.

347

348 **4. Source contributions to BC concentrations, transport and direct radiative** 349 **forcing**

350 **4.1. Source contributions to seasonal mean BC concentrations**

351 Figure 5 shows the simulated spatial distribution of seasonal near-surface BC
352 concentrations originating from the seven tagged source regions in continental China
353 and all other sources from outside China (rest of the world, RW) and Table S3
354 summarizes these source-receptor relationships. It is not surprising that regional
355 emissions largely influence BC concentrations in the same region. For example,
356 emissions of BC from North China give $5.8 \mu\text{g m}^{-3}$ of BC concentrations over North
357 China in DJF, whereas they only account for less than $1.3 \mu\text{g m}^{-3}$ over other regions
358 in China. However, the relatively small amount of BC from upwind source regions can
359 also be a large contributor to receptor regions near the strong sources. BC emissions
360 from North China contribute large amount to concentrations over South, Southwest,
361 Central-West, and Northeast China. BC emissions from South and Southwest China
362 also produce a widespread impact on BC over other neighboring regions. The
363 impacts of BC emitted from the remaining China regions are relatively small both in
364 local and non-local regions due to weak emissions (Fig. 1b). All the sources in China
365 have the largest impact in DJF, resulting from the strong BC emissions in winter,
366 while emissions from outside China have the largest impact on BC over China in
367 MAM due to the seasonal high biomass burning over Southeast Asia and the strong
368 springtime southwesterly winds.

369 Averaged over continental China, emissions of BC from North China produce
370 mean BC concentrations of $0.4\text{--}1.1 \mu\text{g m}^{-3}$, followed by $0.2\text{--}0.4 \mu\text{g m}^{-3}$ from South



371 China and $0.1\text{--}0.2\ \mu\text{g m}^{-3}$ from Southwest China emissions. For emissions over
372 Central-West China, Northeast China, Northwest China, and Tibetan Plateau, their
373 individual impact is less than $0.15\ \mu\text{g m}^{-3}$. In contrast, emissions from outside China
374 result in $0.13\ \mu\text{g m}^{-3}$ of BC concentrations in China in MAM and less than $0.10\ \mu\text{g m}^{-3}$
375 in other seasons. The simulated source contributions to column burden of BC are
376 shown in Figure S3. They present a very similar spatial distribution and seasonal
377 variation to those of near-surface BC concentrations. However, the emissions from
378 outside China have a larger impact on the average column burden of BC over China
379 than on surface concentrations, with a magnitude of $0.5\ \text{mg m}^{-2}$ in MAM, which is as
380 the same as that from sources in North China.

381 Figure 6 shows the spatial distribution of simulated relative contributions to
382 near-surface BC concentrations from sources in the seven regions in continental
383 China and those outside China by season. (The same plots for BC column burden are
384 shown in Figure S4.) For regions with higher emissions, their BC concentrations are
385 dominated by local emissions. In contrast, BC levels, especially column burden of
386 BC, over central and western China with lower emissions are strongly influenced by
387 non-local sources. Emissions from outside China can be the largest contributor to BC
388 over these regions. During DJF, MAM and SON, they contribute more than 70% to
389 both surface concentrations and column burden of BC in Tibetan Plateau, which is
390 important to the climate change due to the large climate efficacy of BC in snow (Qian
391 et al., 2011) and acceleration of snowmelt through elevated BC heat pump
392 mechanism (Lau et al., 2010). BC emissions from outside China also account for a
393 quite significant fraction of surface concentrations over Northwest and Southwest
394 China in MAM, which contribute to poor air quality over these regions.

395 Figure 7 summarizes source attribution for spatially averaged seasonal surface
396 BC concentrations for the seven receptor regions and continental China combined
397 (CN). Over North China, the majority of the BC concentrations are attributed to local
398 emissions in all seasons, with seasonal fractional contributions of 83–93%. Over
399 South China, the seasonal contributions from local emissions are in the range of 64–
400 87%. Emissions from North China account for 30% of BC concentrations over South



401 China in DJF, resulting from the wintertime northwesterly winds (Figure S5a), while
402 emissions from outside China contribute about 10% in MAM due to the strong
403 springtime biomass burning over southeast Asia and southwesterly winds
404 transporting BC from southeast Asia to South China (Figure S5b). Southwest China
405 has a similar level of local influence, with 59–79% of the BC concentration from local
406 emissions, whereas 17% are due to emissions from outside China world transported
407 by westerly winds in MAM.

408 Non-local emissions from Southwest and North China contribute 27–49% of BC
409 concentration in Central-West China. North China emissions play an important role in
410 BC concentrations over Northeast China, with relative contributions in a range of 22–
411 36% in MAM, JJA and SON, while only 12% in DJF, which is associated with
412 northwesterly winds in winter preventing northward transport of BC from North China
413 to Northeast China. Over Northwest China and Tibetan Plateau, 22–40% and 43–
414 76%, respectively, of BC originate from emissions outside China due to the low
415 emissions over the less economically developed western China. For all of continental
416 China as the receptor, the seasonal BC concentrations are largely attributed to the
417 emissions from North China and South China, with relative contributions ranging from
418 43–50% and 18–24%, respectively, followed by contributions from Southwest China
419 (10–13%) and outside China (4–12%).

420 The source region contributions to column burden of BC in each receptor regions
421 in China are shown in Figure S6. In general, impacts on the non-local BC column
422 burden are larger than on surface concentrations because aerosol transport is
423 relatively easier in free-troposphere than in the boundary layer (e.g., Yang et al.,
424 2015). Column burdens of BC averaged over continental China result mainly from
425 emissions in North China and outside China, with relative contributions ranging from
426 31–42% and 14–31%, respectively.

427

428 **4.2. Source contributions during polluted days**

429 Knowing the source attribution of BC during polluted days in China is important
430 for policy makers, which could provide an effective way for the mitigation of poor air



431 quality. Here, the polluted days are simply identified as days with daily concentrations
432 of BC higher than 90th percentile of probability density function in each receptor
433 regions. A total of different 45 days in winter in the 5-year simulation are identified as
434 polluted days for each region in China.

435 Figure 8 shows the DJF composite differences in near-surface BC concentrations
436 and winds at 850 hPa between polluted and normal days for each receptor region,
437 and Figure 9 summarizes the source contributions to the differences. When North
438 China is under the polluted condition, BC concentrations are higher by more than
439 70% compared to DJF average over North China (Fig. 2a), with a maximum increase
440 exceeding $5 \mu\text{g m}^{-3}$. North China local emissions contribute $5.4 \mu\text{g m}^{-3}$ to the
441 averaged increase in BC concentrations over North China during North China
442 polluted days, about 90% of the total increase. In winter, eastern China is dominated
443 by strong northwesterly winds (Figure S4a). The anomalous southerly winds during
444 polluted days (relative to DJF average) over North China prevent the high BC
445 concentrations from being transported to South China, leading to a reduced
446 ventilation and accumulated aerosols in North China.

447 Over South China, BC concentrations increase by up to $2 \mu\text{g m}^{-3}$, in part due to
448 the transport from North China by anomalous northerly winds in the north of South
449 China in South China polluted days. On average, contribution of North China
450 emissions to mean concentrations over South China increases by $1.2 \mu\text{g m}^{-3}$ (48% of
451 total increase) during the South China polluted days.

452 During polluted days in Southwest China, the anomalous northeasterly winds in
453 the east part of Southwest China bring in BC from the highly polluted eastern China,
454 resulting in $1.1 \mu\text{g m}^{-3}$ increase (53% of total increase) in the Southwest China, which
455 is as similar magnitude as the $1.0 \mu\text{g m}^{-3}$ contribution from the Southwest China local
456 emissions.

457 The increase in BC concentrations during polluted days over Central-West China
458 is also largely influenced by the accumulation effect of the anomalous winds over
459 eastern and central China, which also transport BC from Southwest and eastern
460 China into the receptor region.



461 The polluted days in Northeast China are caused by both the accumulation of
462 local emissions due to the reduced prevailing northeasterly winds and anomalous
463 transport of BC from North China.

464 Emissions from outside China could contribute to increases in BC concentrations
465 over Northwest China and Tibetan Plateau during polluted days. However, during
466 wintertime regional polluted days in eastern and central China, the contributions of
467 emissions from outside China do not have a significant influence on the changes in
468 BC concentrations.

469 These results suggest that the transport of aerosols plays an important role in
470 increasing BC concentrations during regional polluted days in eastern and central
471 China. Reductions in local emissions could benefit mitigation of both local and
472 non-local haze in China. Emissions from outside China are not as important to hazy
473 pollution in eastern and central China, where haze episodes occur frequently in winter
474 due to relatively high anthropogenic aerosol emissions and abnormal meteorological
475 conditions (Sun et al., 2014; R. H. Zhang et al., 2014; Yang et al., 2016). Note that, in
476 this study, we only focus on the source-receptor relationships related to the wind
477 anomalies during polluted days. In addition to winds, changes in other meteorological
478 fields, such as precipitation, temperature, humidity, and planetary boundary layer
479 height, could also influence the contributions of local aerosols between polluted and
480 normal days. Although the BC emissions used in the simulation include a seasonal
481 variability that could cause some variations in simulated concentrations, the monthly
482 variability in DJF of BC emissions is less than 4% over China, which is negligible
483 compared to the differences in concentrations between polluted and normal days.

484

485 **4.3. Source contributions to trans-boundary and trans-Pacific transport**

486 Considering the large contributions of emissions from Southeast Asia to MAM BC
487 concentrations in the southwest China (Figure 6) and the large outflow of aerosols
488 from East Asia in springtime (Yu et al., 2008), it is valuable to examine the inflow and
489 outflow of BC in China. Figures S7a and S7b show the vertical distribution of source
490 contributions of emissions from outside China to BC concentrations averaged over



491 75°–120°E and 25°–35°N, respectively, around the south boundary of continental
492 China in MAM. High concentrations of BC originating from Southeast Asia are lifted to
493 the free atmosphere in the south slope of Tibetan Plateau. Then westerly winds
494 transport these BC particles to Southwest China and South China in both low- and
495 mid-troposphere. Figures S7c and S7d present the contributions of emissions from
496 China to BC concentrations averaged over 120°–135°E and 20°–50°N, respectively,
497 around the east boundary of continental China. In MAM, the northward meridional
498 winds over 25°–35°N and the southward meridional winds over 40°–50°N lead to the
499 accumulation of BC in the lower atmosphere in eastern China. Westerly winds then
500 transport these BC out of China mostly under 500 hPa.

501 Figure 10 shows the spatial distribution of column burden and surface
502 concentrations of BC resulting from emissions in and outside China in MAM. Column
503 burden is used to represent the outflow in this study following previous studies (Chin
504 et al., 2007; Hadley et al., 2007). There are strong outflows across the Pacific Ocean
505 originating from emissions both in and outside China. Emissions from China
506 contribute 0.19 mg m⁻² (or 53%) of MAM mean BC along 150°E averaged over 20°–
507 60°N, whereas emissions outside China contribute 0.17 mg m⁻² (or 47%). It suggests
508 that both emissions from China and outside China are important for the outflow from
509 East Asia. The yearly contribution from emissions from China in this study is 58%,
510 similar to 61% in Matsui et al. (2013) calculated based on eastward BC mass flux
511 using WRF-CMAQ model with INTEX-B missions. Averaged over western United
512 States (125°–105°W, 30°–50°N), emissions from China account for 7% of
513 near-surface BC concentrations and 25% in column burden in MAM, indicating that
514 emissions from China could have a significant impact on air quality in western United
515 States. More than half of the China contribution to BC over western United States
516 originates from eastern China (i.e., the tagged North and South China).

517

518 4.4. Source contributions to direct radiative forcing

519 The high concentrations of BC in China could also have a significant impact on
520 the climate system through atmospheric heating or direct radiative forcing. As shown



521 in Figure 11, the annual mean direct radiative forcing (DRF) of BC at TOA is as high
522 as $3\text{--}4\text{ W m}^{-2}$ at some locations. Similar to the source attributions of BC
523 concentrations (Figure 5) and burden (Figure S3), regional sources contribute the
524 largest to DRF over the respective local regions. Among all the source regions in
525 China, emissions from North, South, and Southwest China contribute the largest to
526 local DRF of BC, with maximum DRF in a range of $3\text{--}5$, $2\text{--}3$, and $3\text{--}5\text{ W m}^{-2}$,
527 respectively. Other sources regions in China have relatively low contributions, with
528 maximum values less than 2 W m^{-2} . Emissions outside China lead to $1\text{--}2\text{ W m}^{-2}$ of
529 DRF of BC over South, Southwest, Northwest China and Tibetan Plateau, and $0.2\text{--}1$
530 W m^{-2} over other parts of China, an effect that is quite widespread.

531 The total DRF of BC averaged over continental China simulated in this study is
532 2.27 W m^{-2} , larger than $0.75\text{--}1.46\text{ W m}^{-2}$ in previous studies (Wu et al., 2008; Zhuang
533 et al., 2011; Zhuang et al., 2013; Li et al., 2016), probably due to the different
534 emissions in the time periods of study. The total BC emissions averaged over
535 continental China were 1005 Gg C yr^{-1} for years 1993–2003 in Wu et al. (2008), 1811
536 Gg C yr^{-1} for year 2006 in Zhuang et al. (2011, 2013), and 1840 Gg C yr^{-1} for year
537 2010 in Li et al. (2016), whereas 2497 Gg C yr^{-1} for year 2010–2014 used in this
538 study. Emissions outside China have the largest contributions to DRF of BC in China
539 compared to any of the individual source regions in China, with an averaged
540 contribution of 0.77 W m^{-2} (34%). This fractional contribution from emissions outside
541 China is larger than 25% in Li et al. (2016), however we use different emissions,
542 model and meteorology. Emissions from North China result in 0.56 W m^{-2} (25%) of
543 DRF of BC over China, followed by 0.33 W m^{-2} (15%) and 0.31 W m^{-2} (14%) from
544 South China and Southwest China, respectively. Emissions from Central-West,
545 Northeast, Northwest China, and Tibetan Plateau taken together account for 0.30 W
546 m^{-2} (13%) of DRF of BC over China.

547 Figure 12a shows the seasonal mean DRF of BC averaged over China as a
548 function of regional BC emissions. Because of high emissions, DRF of BC emitted
549 from North China is the largest in all seasons, with values in a range of $0.5\text{--}0.7\text{ W m}^{-2}$
550 averaged over China, followed by $0.2\text{--}0.5\text{ W m}^{-2}$ from South and Southwest China.



551 BC from the other tagged regions in China contribute less than 0.2 W m^{-2} in all
552 seasons. In general, BC DRF in each season is proportional to its emission rate.

553 Figure 12b presents the seasonal DRF efficiency of BC emitted from the tagged
554 regions and Table S4 summarizes these efficiencies. The variability of DRF efficiency
555 for forcing over China is determined by several factors, such as incoming solar
556 radiation (location of source regions), BC column burden and vertical distribution, and
557 transport out of the region. The China DRF efficiency is largest in western China (NW
558 and TP). This spatial pattern was also found by Henze et al. (2012). It can be
559 explained by the increase of multiple scattering effects and attenuation of the
560 transmitted radiation for large AOD (García et al., 2012). The Northeast China region
561 has a low China DRF efficiency due to transport eastward outside of China. The
562 remaining central and southern China regions have China DRF efficiencies that are
563 fairly consistent, varying by 20-30% about the average.

564 DRF efficiencies of BC from most regions have higher values in JJA and lower
565 values in DJF. This is primarily due to more incoming solar radiation in summer.
566 Insolation is the largest over Northwest China in JJA, together with less precipitation
567 than other regions, resulting in large DRF efficiency there. Global BC DRF efficiency,
568 particularly the annual average, is fairly similar for central, southern, and eastern
569 China regions (Fig. 12c, d). Global efficiency is still much higher for the western
570 regions.

571 BC emission reductions may impact mitigation of climate change and improve air
572 quality. To compare the relative importance of climate and air quality effects of BC
573 from different regions in China, Fig. 13 shows the near-surface concentration and
574 column burden efficiency of BC over China and globally and Table S5 summarizes
575 these efficiencies. For near-surface concentration (Fig. 13a and 13b), the efficiencies
576 are largest in DJF and lowest in JJA, in contrast to the DRF efficiencies, resulting
577 from the less precipitation and wet deposition of aerosols in winter. Unlike the DRF
578 efficiencies, the near-surface concentration efficiencies over eastern China are
579 similar and even larger than those for central and western China. These results



580 suggest that reduction in BC emissions in eastern China could benefit more on the
581 regional air quality in China, especially in winter haze season.

582 The relative distributions of column burden efficiencies (Fig. 13c and 13d) are
583 similar to the DRF efficiencies for the major emitting region in China, indicating that
584 aerosol lifetime in atmosphere drives DRF that influences regional and global climate.
585 The western regions (NW and TP), as expected, have a higher forcing per unit
586 column burden.

587

588

589 **5. Conclusions**

590 In this study, the Community Earth System Model (CESM) with a source-tagging
591 technique is used to quantify the contributions of BC emitted from seven regions in
592 continental China, including North China (NC), South China (SC), Southwest China
593 (SW), Central-West China (CW), Northeast China (NE), Northwest China (NW), and
594 Tibetan Plateau (TP), and sources outside China (RW) to concentrations, haze
595 formation, trans-boundary and trans-Pacific transport, and direct radiative forcing
596 (DRF) of BC in China. The anthropogenic emissions of BC for years 2009-2014 used
597 in this study were developed for the Coupled Model Intercomparison Project Phase 6
598 (CMIP6) from the Community Emissions Data System (CEDS). The annual total
599 emission of BC from continental China is 2497 Gg C averaged over years 2010–2014.
600 The model captures well the spatial distribution and seasonal variation in China.
601 AAOD compares well with measurements, which are largely located in central and
602 eastern China. Surface BC concentrations are underestimated by 53% compared to
603 point observations.

604 The individual source regions are the largest contributors to their local BC
605 concentration levels. Over North China where the air quality is often poor, about 90%
606 of near-surface BC concentration is contributed by local emissions. However, some
607 source regions also impact BC in neighboring regions. Due to the seasonal variability
608 of winds and emission rates, emissions from North China account for 30% of
609 near-surface BC concentrations over South China in DJF



610 (December-January-February), while emissions from outside China contribute about
611 10% in MAM (March-April-May). Over Southwest China, 17% of BC in MAM comes
612 from sources outside China. Southwest and North China emissions contribute largely
613 to BC in Central-West China. North China emissions have a contribution in a range of
614 12–36% to BC concentrations in Northeast China. Over Northwest China and Tibetan
615 Plateau, more than 20% and 40% of BC, respectively, originates from emissions
616 outside China. These indicate that, for regions with high emissions, their BC
617 concentrations are dominated by local emissions. In contrast, BC levels over central
618 and western China with lower emissions are more strongly influenced by non-local
619 emissions. For all continental China as a whole, seasonal BC concentrations are
620 largely due to emissions from North and South China, with relative contributions
621 ranging from 43–50% and 18–24%, respectively, followed by contributions from
622 Southwest (10–13%) and outside China (4–12%).

623 Emissions from non-local sources together with abnormal winds are one of the
624 important factors contributing to high winter time pollution events in China. Over
625 South China, about 50% of the increase in BC concentrations during high pollution
626 conditions results from North China emissions. The increases in BC concentrations
627 during polluted days over Southwest, Central-West and Northeast China are strongly
628 influenced by emissions from eastern China. Emissions from outside China could
629 contribute significantly to increases in BC concentrations over Northwest China and
630 Tibetan Plateau during their polluted days. However, emissions from outside China
631 do not have a significant contribution to haze in eastern and central China,
632 suggesting that reduction in emissions within China would be needed to mitigate both
633 local and non-local BC concentrations under high-polluted conditions.

634 Emissions from regions in and outside China both account for about half of BC
635 outflow from East Asia, suggesting that emissions from China and other regions are
636 equally important for the BC outflow from East Asia. Through long-range transport,
637 emissions from China result in 7% of near-surface BC concentration and 25% in
638 column burden over western United States in MAM, indicating that emissions from
639 China are could have an impact on air quality in western United States.



640 The total DRF of BC averaged over continental China simulated in this study is
641 2.27 W m^{-2} . Among the tagged regions, emissions outside China have the largest
642 single contribution to DRF of BC in China, with an average contribution of 34%,
643 followed by 25%, 15%, and 14% due to emissions from North, South and Southwest
644 China, respectively. DRF efficiencies over eastern China are small compared to
645 central and western China in all seasons. It can be explained by the increase of
646 multiple scattering effects and attenuation of the transmitted radiation for large AOD.
647 For near-surface concentration, the efficiencies are largest in DJF and lowest in JJA,
648 and efficiencies over eastern China are similar and even larger than central and
649 western China. These suggest that reduction in BC emissions over eastern China
650 could benefit more on the regional air quality in China, especially in winter haze
651 season.

652 Note that the model largely underestimates BC concentrations over China,
653 compared to the observation, which has also been reported in many previous studies
654 using different models and different emission inventories (e.g., Liu et al., 2012; Fu et
655 al., 2012; H. Wang et al., 2013; R. Wang et al., 2014; Li et al., 2016). One possible
656 reason is that in situ measurements are point observations, while the model does not
657 treat the subgrid variability of aerosols and assumes aerosols are uniformly
658 distributed over the grid cell. R. Wang et al. (2014) found a reduction of negative bias
659 (from -88% to -35%) in the modeled surface BC concentrations when using
660 high-resolution emissions and modeling at $0.5^\circ \times 0.7^\circ$ resolution. They find, however,
661 that modeling over the North China Plain at an even higher resolution of 0.1° , further
662 reduced the surface concentration bias there from 29% to 8%. This result indicates
663 that the siting of observational stations can result in an artificial bias when comparing
664 with relatively coarse model results. Further investigation of this siting/resolution bias
665 is warranted, including investigation if this type of bias might extend, presumably to a
666 lesser extent, also to AAOD measurements.

667 Further reasons that could contribute to this bias are emission underestimation or
668 inaccurate aerosol processes in the model. Given that the differences between
669 modeled observed AAOD over eastern China are relatively small (-18%), we



670 conclude that, given current evidence, the total amount of atmospheric BC in these
671 simulations is reasonable at least in this sub-region.

672 Over eastern China, the BC concentrations are dominated by local emissions in
673 this study, with local contribution of 64–93%. The underestimation of simulated BC
674 concentrations over eastern China is more likely due to either underestimation of
675 local emissions, too much aerosol removal within these regions, or resolution bias
676 between observations and model grids. Over western China, 22–76% of the BC
677 originates from emissions outside China. Thus biases of simulated BC concentrations
678 could also come from underestimation of emissions outside China and or too much
679 removal of BC during long-range transport. Satellite data are a promising method to
680 validate modeling and emissions inventories, given that they do not depend on the
681 location of observing stations, providing more uniform spatial coverage. A
682 comparison of modeled AAOD and satellite aerosol index (AI) provides an indication
683 that the modeled burden in western China is underestimated, although the role of
684 dust needs to be better characterized.

685

686

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972 **Figure Captions**

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974 **Figure 1.** (a) Spatial distribution of annual mean total emissions (anthropogenic plus
975 biomass burning, units: $\text{g C m}^{-2} \text{ yr}^{-1}$) of black carbon (BC) averaged over 2010–2014.
976 The geographical BC source regions are selected as North China (NC, 109°E–east
977 boundary, 30°–41°N), South China (SC, 109°E–east boundary, south boundary–
978 30°N), Southwest China (SW, 100°–109°N, south boundary–32°N), Central-West
979 China (CW, 100°–109°N, 32°N–north boundary), Northeast China (NE, 109°E–east
980 boundary, 41°N–north boundary), Northwest China (NW, west boundary–100°E,
981 36°N–north boundary), and Tibetan Plateau (TP, west boundary–100°E, south
982 boundary–36°N) in China and regions outside of China (RW, rest of the world). (b)
983 Seasonal mean total emissions (units: Gg C , $\text{Gg} = 10^9\text{g}$) of BC from the seven BC
984 source regions in China.

985

986 **Figure 2.** Simulated seasonal mean near-surface concentrations (left, units: $\mu\text{g m}^{-3}$)
987 and column burden (right, units: mg m^{-2}) of BC in December-January-February (DJF),
988 March-April-May (MAM), June-July-August (JJA), and
989 September-October-November (SON).

990

991 **Figure 3.** Comparisons of observed and modeled seasonal mean (a) near-surface
992 concentrations (units: $\mu\text{g m}^{-3}$) and (b) aerosol absorption optical depth (AAOD) of BC
993 in China. Solid lines mark the 1:1 ratios and dashed lines mark the 1:3 and 3:1 ratios.
994 Observed BC concentrations were taken between 2006 and 2007 at 14 sites of the
995 China Meteorological Administration (CMA) Atmosphere Watch Network (CAWNET)
996 (Zhang et al., 2012). Observed AAOD of BC are obtained by removing dust AAOD
997 from total AAOD at 10 sites of the Aerosol Robotic Network (AERONET) (Holben et
998 al., 2001), following Bond et al. (2013). The observed AAOD are averaged over years
999 of 2005–2014 with data available. Correlation coefficient (R) and normalized mean
1000 bias (NMB) between observation and simulation are shown on top left of each panel.



1001 $NMB = 100\% \times \sum (M_i - O_i) / \sum O_i$, where M_i and O_i are the modeled and observed
1002 values at site i , respectively. Site locations are shown in Figure S1a.

1003

1004 **Figure 4.** Spatial distribution of seasonal mean AAOD of total aerosols (left) and
1005 Aerosol Index (AI) derived from Ozone Monitoring Instrument (OMI) measurements
1006 over years of 2010–2014 (right).

1007

1008 **Figure 5.** Spatial distribution of seasonal mean near-surface concentrations of BC
1009 ($\mu\text{g m}^{-3}$) originating from the seven source regions in China (NC, SC, SW, CW, NE,
1010 NW, and TP), marked with black outlines, and sources outside China (RW).
1011 Regionally averaged BC in China contributed by individual source regions is shown at
1012 the bottom right of each panel.

1013

1014 **Figure 6.** Spatial distribution of relative contributions (%) to seasonal mean
1015 near-surface BC concentrations from each of the tagged source regions.

1016

1017 **Figure 7.** Relative contributions (%) from the tagged source regions (denoted by
1018 colors) to regional mean surface concentrations of BC over seven receptor regions in
1019 China (NC, SC, SW, CW, NE, NW, and TP) and China (seven regions combined, CN)
1020 in different seasons. The receptor regions are marked on the horizontal axis in each
1021 panel.

1022

1023 **Figure 8.** Composite differences in winds at 850 hPa (m s^{-1}) and near-surface BC
1024 concentrations ($\mu\text{g m}^{-3}$) between polluted and normal days in DJF.

1025

1026 **Figure 9.** Composite differences in surface BC concentrations ($\mu\text{g m}^{-3}$) averaged
1027 over receptor regions (marked on the horizontal axis) over eastern and central China
1028 between polluted and normal days in DJF originating from individual sources regions
1029 (bars in each column).

1030



1031 **Figure 10.** Spatial distribution of (a, b) column burden (mg m^{-2}) and (c, d)
1032 near-surface concentrations ($\mu\text{g m}^{-3}$) of BC originating from total emissions inside
1033 (CN) and outside China (RW), respectively, in March-April-May (MAM). The black
1034 solid lines over western (150°E , $20^\circ\text{--}60^\circ\text{N}$) Pacific in panel (a) mark the
1035 cross-sections used to quantify outflow of BC from East Asia. The box over western
1036 United States ($125^\circ\text{--}105^\circ\text{W}$, $30^\circ\text{--}50^\circ\text{N}$) in panel (c) is used to quantify BC
1037 concentrations attributed to sources from China.

1038

1039 **Figure 11.** Spatial distribution of annual mean direct radiative forcing of BC (W m^{-2}) at
1040 the top of the atmosphere originating from the tagged BC source regions in China
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1042 averaged forcing in China contributed by individual source regions is shown at the
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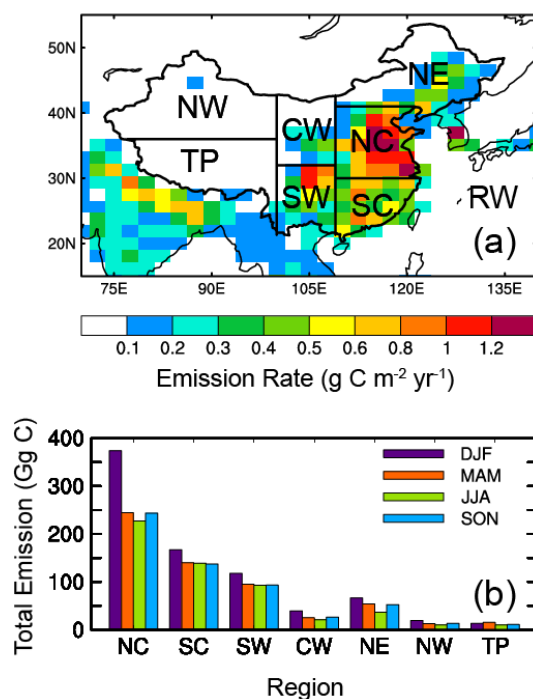
1045 **Figure 12.** (a, c) BC seasonal DRF averaged over China as a function of BC
1046 emission fraction (the ratio of regional emission to the total emission over China and
1047 global, respectively, unit: %) for each of the tagged regions. (b, d) Seasonal DRF
1048 efficiency of BC ($\text{W m}^{-2} \text{Tg}^{-1}$) for each of the tagged source regions over China and
1049 globally, respectively. The efficiency is defined as the DRF divided by the
1050 corresponding scaled annual emission (seasonal emission multiplied by 4). Error bars
1051 indicate $1\text{-}\sigma$ of mean values during years 2010–2014.

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1053 **Figure 13.** Seasonal (a, b) near-surface concentration ($\mu\text{g m}^{-3} \text{Tg}^{-1}$) and (c, d) column
1054 burden ($\text{mg m}^{-2} \text{Tg}^{-1}$) efficiency of BC for each of the tagged source regions over
1055 China and globally, respectively.

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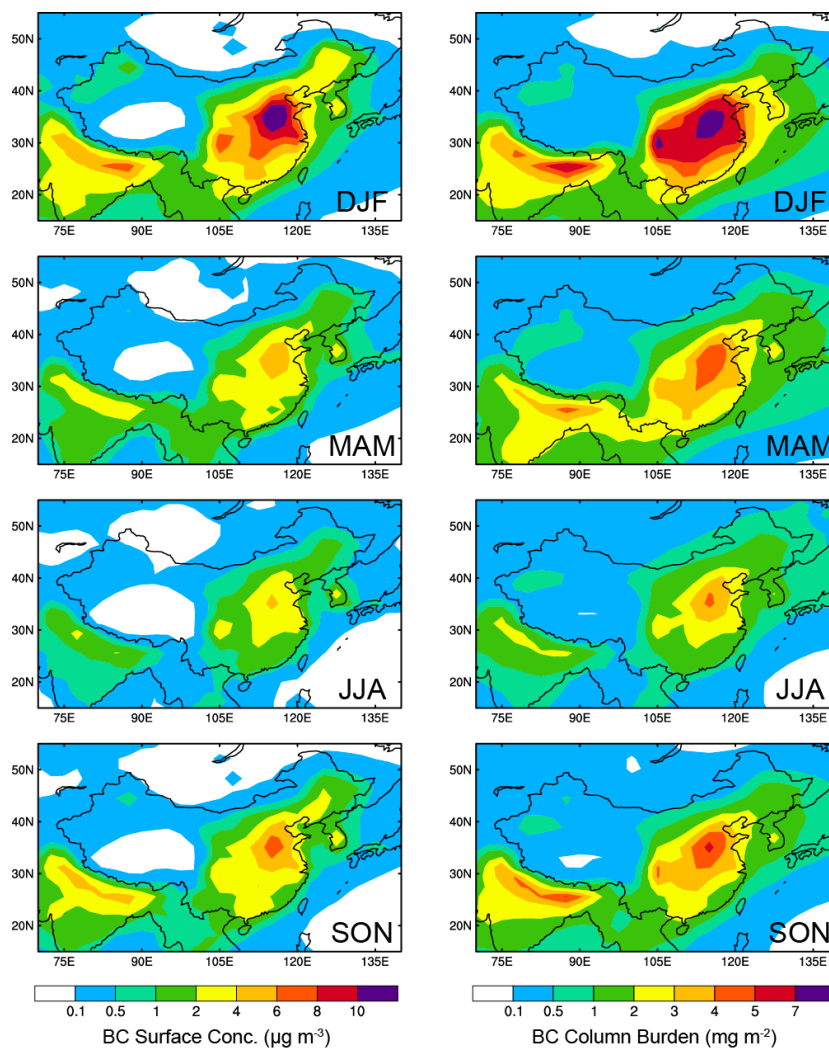


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1065 China (CW, 100° – 109°E , 32°N –north boundary), Northeast China (NE, 109°E –east
1066 boundary, 41°N –north boundary), Northwest China (NW, west boundary– 100°E ,
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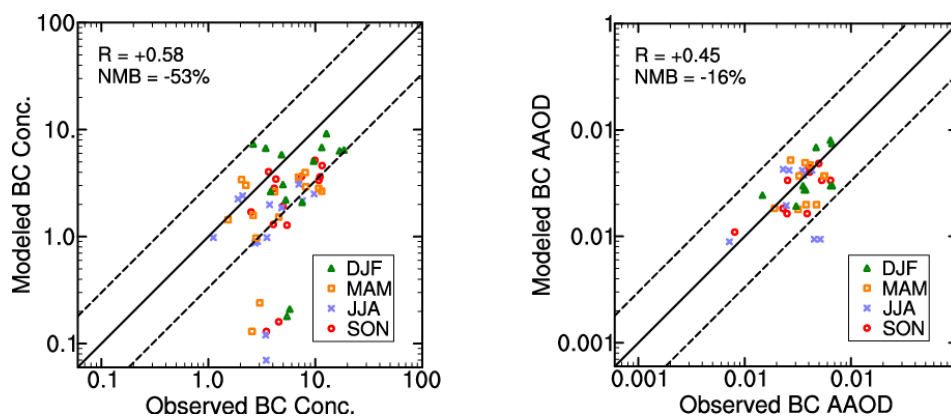
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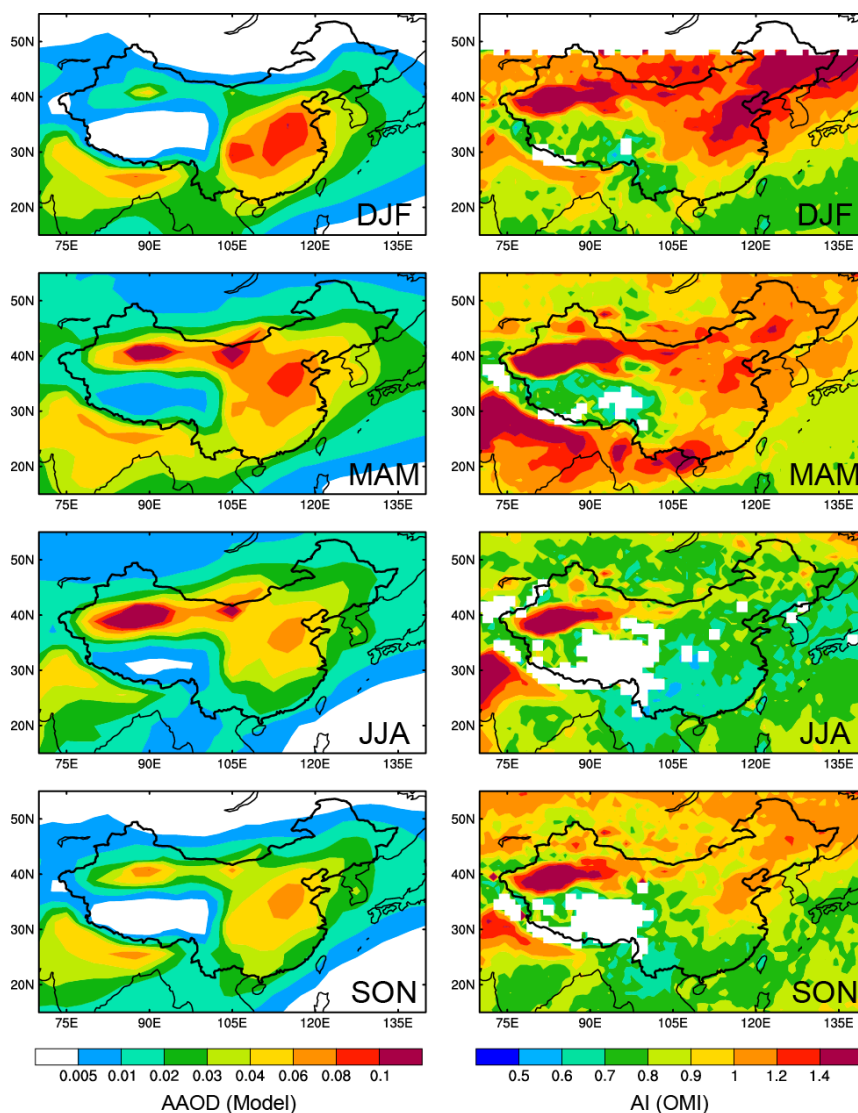
1074 **Figure 2.** Simulated seasonal mean near-surface concentrations (left, units: $\mu\text{g m}^{-3}$)
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1079

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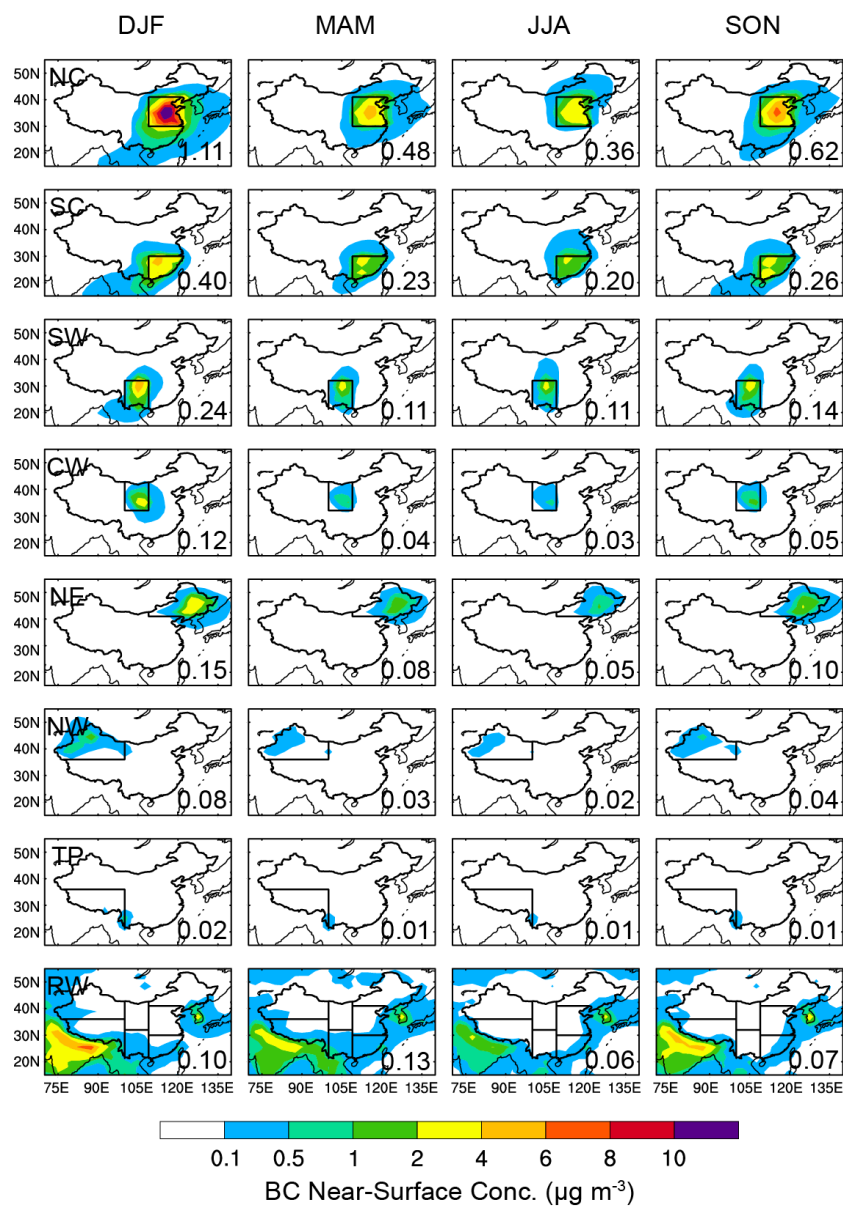
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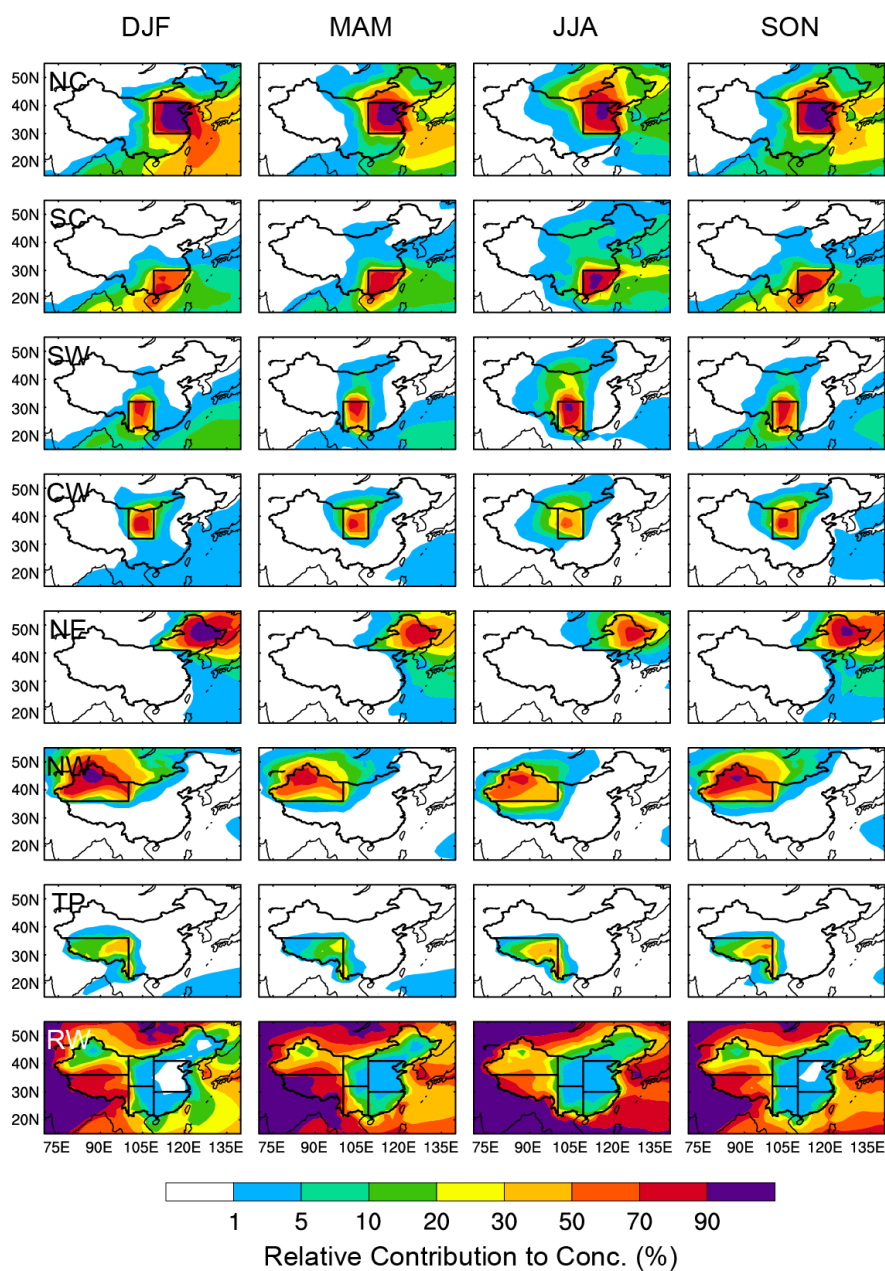


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1102 **Figure 5.** Spatial distribution of seasonal mean near-surface concentrations of BC
 1103 ($\mu\text{g m}^{-3}$) originating from the seven source regions in China (NC, SC, SW, CW, NE,
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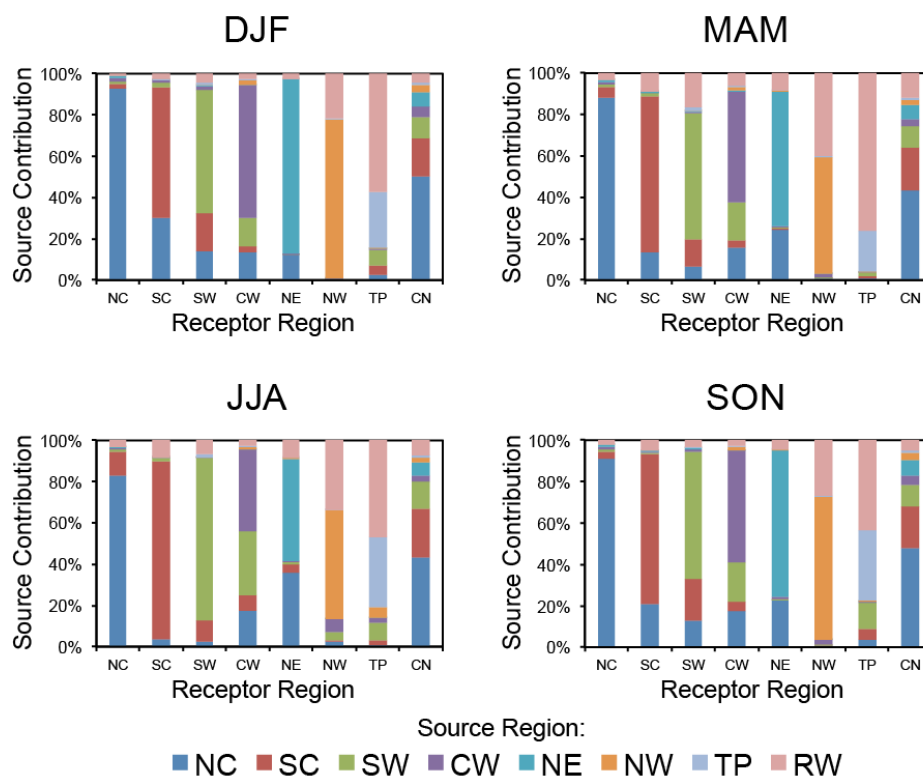
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1109 **Figure 6.** Spatial distribution of relative contributions (%) to seasonal mean

1110 near-surface BC concentrations from each of the tagged source regions.

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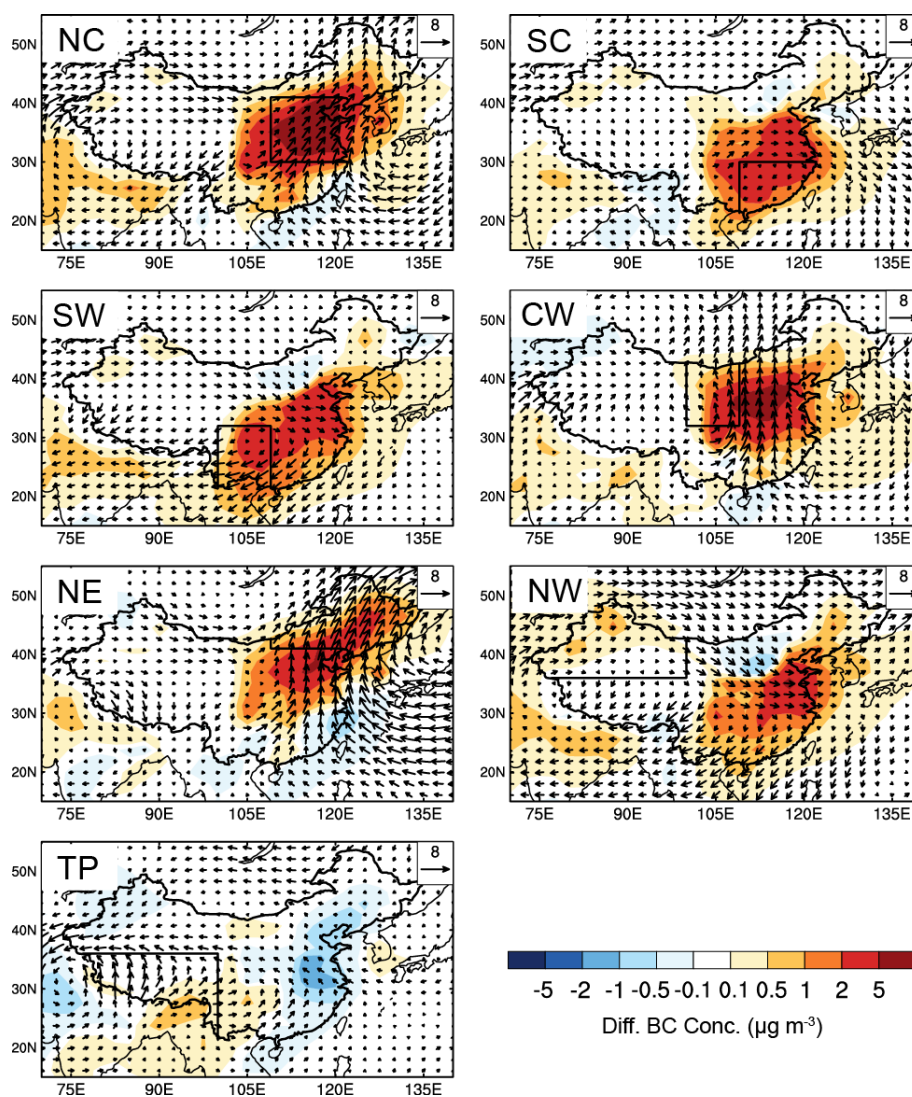
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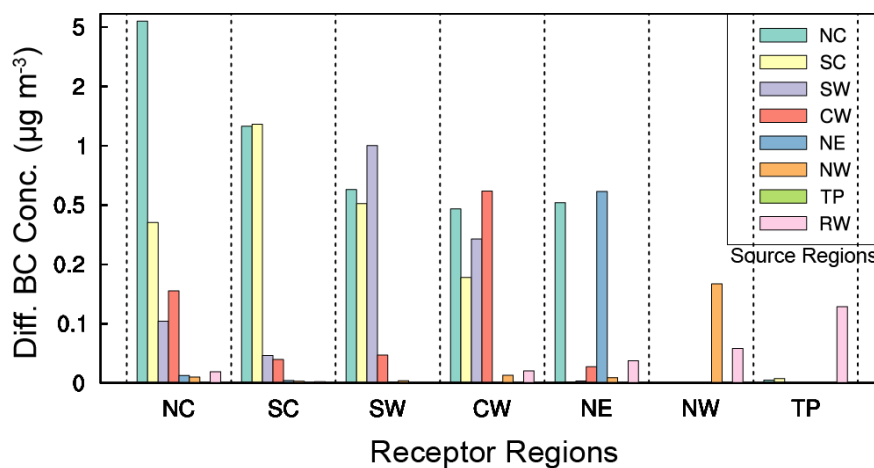
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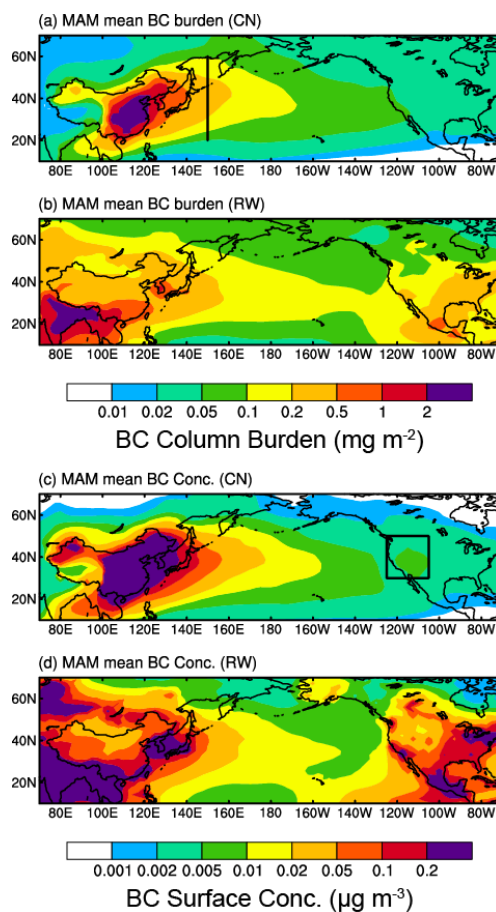


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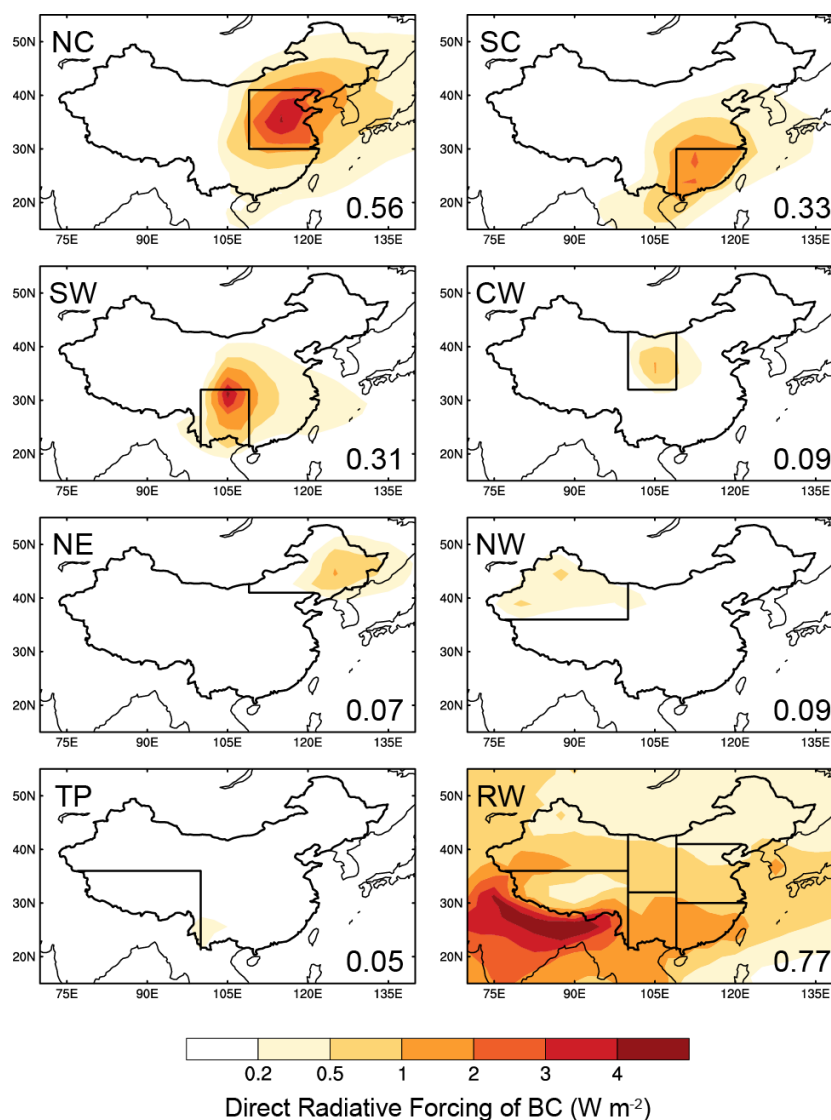
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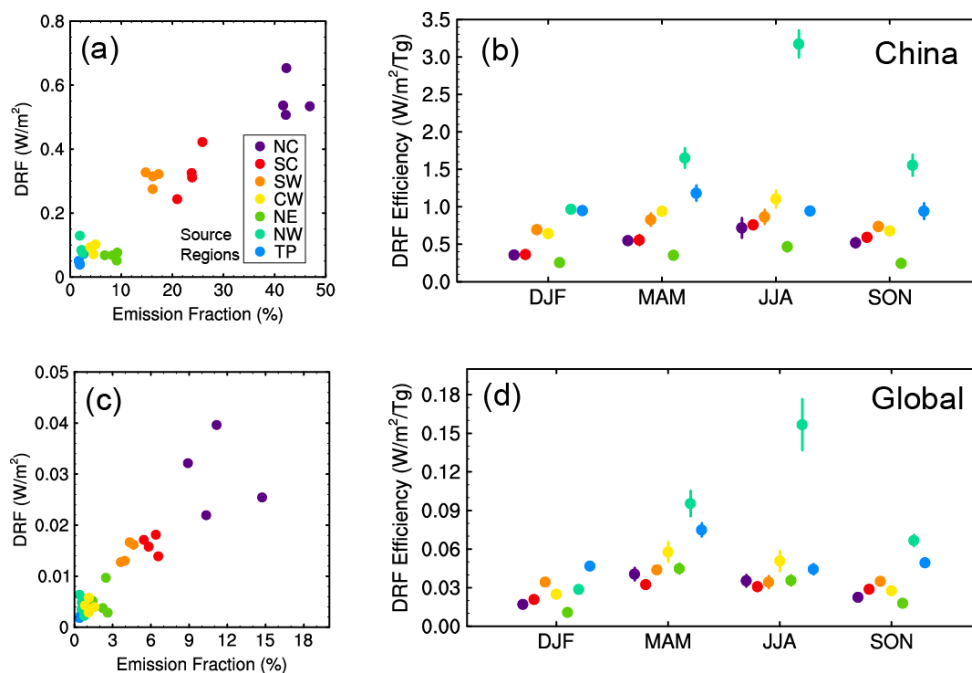
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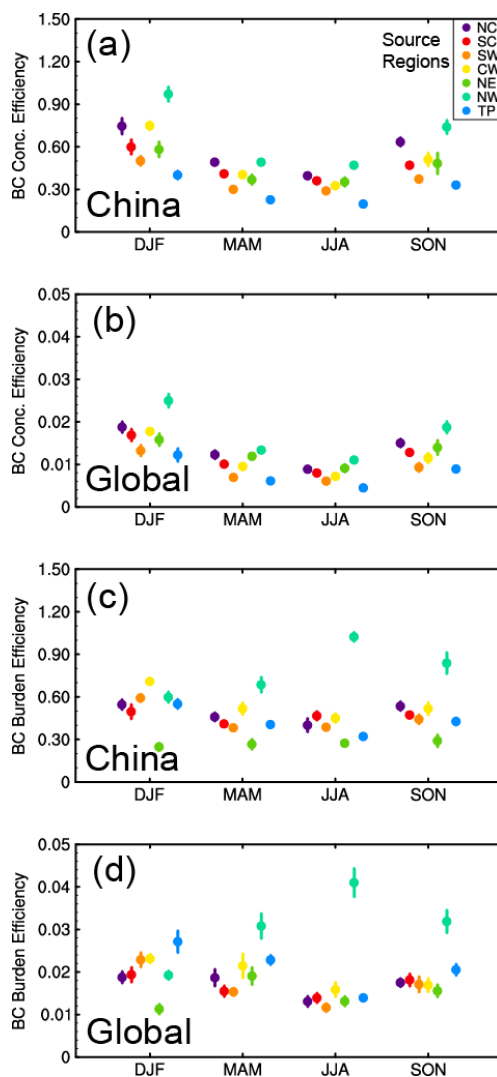


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