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2 **On the simultaneous deployment of two**
3 **single particle mass spectrometers at an**
4 **urban background and a road side site**
5 **during SAPUSS**

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1 **Abstract**

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3 The Aerosol Time-of-Flight Mass Spectrometer (ATOFMS) provides size resolved
4 information on the chemical composition of single particles with high time resolution.
5 Within SAPUSS (Solving Aerosol Problems by Using Synergistic Strategies),
6 continuous ATOFMS measurements of ambient particles were made simultaneously
7 at two urban locations: urban background (UB) site and road side (RS) site in the city
8 of Barcelona (Spain) from 17th September to 18th October 2010. Two different
9 instrumental configurations were used: ATOFMS (TSI 3800) with a converging
10 nozzle inlet (high efficiency at about 800-2000nm) at the UB site and ATOFMS (TSI
11 3800-100) with an aerodynamic lens inlet (high efficiency at about 300-700nm) at the
12 RS site. This is the first time, to our knowledge, that two ATOFMS instruments have
13 been deployed in the same field study. The different instrument configurations had an
14 impact on the observed particle types at the two sites. Nevertheless, ten particle
15 types were detected at both locations, including local and regional elemental carbon
16 (22.7-58.9% of total particles), fresh and aged sea salt (1.0-14.6%), local and
17 regional nitrate -containing aerosols (3-11.6%), local lead-containing metallic
18 particles (0.1-0.2%) and transported Fe-nitrate particles (0.8-2.5%). The ATOFMS at
19 the UB also characterised four particle types: calcium-containing dust (0.9%),
20 Saharan dust (1.3%), vanadium-containing particles (0.9%) and vegetative debris
21 (1.7%). By contrast, the high statistical counts of fine particles detected at the RS
22 allowed identification of eight particle types. Four of these contained amines of
23 primary and secondary origin. Aminium salts were found related to coarse sulphate
24 rich particle types, suggesting heterogeneous reaction mechanisms for their



1 formation. The other four particle types mainly containing organic carbon were found
2 spiking at different types of the day, showing a complex single particle mixing state
3 relationship between organic carbon and nitrate. This ATOFMS study clearly shows
4 that the composition of atmospheric fine particles in Barcelona, and likely other
5 Mediterranean urban areas, is complex, with a wide range of local and regional
6 sources combining with chemical processing to produce at least twenty-two different
7 particle types exhibiting different temporal behaviour. The advantage of using two
8 ATOFMS instruments is also demonstrated, with the nozzle-skimmer configuration
9 enabling detection of coarse dust particles and the aerodynamic lens configuration
10 allowing better identification of particles rich in organic carbon and amines.

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3 **1. Introduction**

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5 A substantial number of studies have shown a relationship between measures of
6 particulate air pollution and a variety of adverse health indicators (WHO, 2004).
7 Formulation of cost-effective air pollution control policies depends upon a sound
8 knowledge of source contributions to ambient concentrations. Only with such
9 knowledge can realistic cost-benefit evaluations be conducted. The major sources of
10 ambient particles in most urban areas are primary emissions from road traffic and
11 other fuel combustion, secondary particles arising from condensation or chemical
12 processing, and resuspension of soils and road dusts (AQEG, 2005; Harrison et al.,
13 2012). Marine aerosol can also contribute in coastal locations and the interactions of
14 anthropogenic trace gases with natural aerosol (i.e. dust, sea salt) can also have
15 significant effects on aerosol composition (Abbatt et al., 2012).

16 Measurement of particle composition by on-line mass spectrometry has developed
17 extensively over the last two decades and is currently the fastest growing area of
18 atmospheric aerosol research (Laskin et al., 2012). The aerosol time-of-flight mass
19 spectrometer (ATOFMS) has been used in many previous field studies to determine
20 the chemical constituents of atmospheric aerosols (Pratt and Prather, 2012). It can
21 identify both refractory and non-refractory species in single particles and can provide
22 size-resolved information on particle sources and atmospheric processing at high
23 time resolution (Prather and Pratt, 2012; Laskin et al., 2012;). The ATOFMS has
24 been used in a number of recent field studies in urban areas of Europe (Dall'Osto



1 and Harrison, 2012; Healey et al., 2013) to identify and characterise particles from a
2 diverse range of anthropogenic sources including traffic, solid fuel burning, industry,
3 soil and road dust, marine aerosol and secondary aerosol formation processes.
4 However, it is worthy of mention that ATOFMS source apportionment capabilities are
5 limited by the difficulties in quantification of its outputs (Reilly et al., 2000; Schoolcraft
6 et al., 2001). However, single-particle analysis is an important analytical tool that
7 allows us to determine how the myriad chemical constituents are distributed between
8 individual particles (mixing state, Pratt and Prather, 2012). The ATOFMS has often
9 reported a number of particle types, which at times are difficult to associate with a
10 specific aerosol source (Pastor et al., 2003; Dall’Osto and Harrison 2006; 2012).

11 The objective of the present manuscript is to report a detailed analysis of the
12 ATOFMS particle types detected during a field measurement campaign carried out in
13 Barcelona, Spain, as part of the SAPUSS (Solving Aerosol Problems by Using
14 Synergistic Strategies) project . The ATOFMS cannot provide quantitative aerosol
15 mass loading concentrations, but its strength relies in the fact that it can monitor in
16 real time variations in the single particle composition. In other words, small variations
17 in the particle mixing state results in a single particle mass spectra. As a result, a
18 number of atmospheric processes and aerosol sources can be monitored in real
19 time. In this paper we discuss, not only information on the mass spectra, but also
20 diurnal trends persisting over four weeks are discussed. Further information on the
21 intensive field campaign can be found in the presenting overview paper found in this
22 ACP special issue (Dall’Osto et al., 2012a). Two different ATOFMS instruments
23 were deployed during the four-week measurement period - one at a Road Side (RS)
24 site and the other at an Urban Background (UB) site. This is the first time an
25 ATOFMS has been deployed in Spain and, to the best of our knowledge, it is also the



1 first time (worldwide) that two ATOFMS instruments have been deployed
2 simultaneously in the same field campaign. The similarities and differences in particle
3 types detected at both sites is described in detail and attributed to a range of local
4 and regional sources, as well as to different chemical and physical processes.

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6 **2. Methods**

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8 **2.1 Location**

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10 The SAPUSS field measurement campaign involved a large variety of
11 instrumentation deployed simultaneously at a number of monitoring sites in
12 Barcelona (Spain), between 17th September and 18th October 2010 (local time,
13 UTC+2) (Dall'Osto et al., 2012). ATOFMS measurements were made at the two main
14 SAPUSS supersites:

- 15 • Road side (RS) site was situated in a car park next to a major road (Carrer
16 Urgell). The road, which crosses the city from South East to North West, is a
17 street canyon composed of a two-way cycling path and a one-way four lane
18 vehicle road. Vehicle intensity for the month of measurements was about 17,000
19 vehicles per day.
- 20 • Urban Background (UB) site was situated in a small park at the North Western
21 periphery of the city centre. A main road (Avenida Diagonal, 127,000
22 vehicles/day) is located about 500 m away from the site.



1 The two sites were about 2 km from each other (Dall'Osto et al., 2012). While the UB
2 site was open to wind from all directions, the wind flow and turbulence at the RS site
3 were affected by the nearby street canyons and vehicular traffic.

4 **2.2 Instrumentation**

5 The mass spectrometers were housed in air-conditioned trailers at both sites.
6 Sampling was performed ca. 4 m above ground using a quarter-inch internal
7 diameter stainless steel tube fitted with a PM_{2.5} cyclone. The sample air was dried
8 before arriving at the instruments. The two ATOFMS instruments used in this study
9 had different configurations. The instrument deployed at the RS site was an
10 ATOFMS TSI model 3800-100, in which particles are sampled through an orifice and
11 accelerated through an aerodynamic lens to the sizing region of the instrument (Su et
12 al. 2004). By contrast, the instrument at the UB site was an ATOFMS TSI Model
13 3800 that utilized a converging nozzle inlet (Gard et al. 1997). Both instruments
14 provide the aerodynamic diameter of particles sizes between about 100 nm and 3 μm
15 by calculating their time of flight between two orthogonally positioned continuous
16 wave lasers ($\lambda = 532$ nm). However, the transmission efficiencies of the two
17 instruments are quite different. While the aerodynamic lens affords a much higher
18 transmission efficiency for particles with diameters less than about 1 μm, its
19 performance for larger particles is not as good as the converging nozzle inlet.
20 Following the sampling and sizing, particles are transferred to the mass spectrometry
21 region where a pulsed laser ($\lambda = 266$ nm, about 1 mJ pulse⁻¹) desorbs and ionizes
22 material within the particle in the centre of the ion source of a bipolar reflectron ToF-
23 MS. Thus, positive and negative ion mass spectra of a single particle are obtained.
24 Overall, during the SAPUSS field study the ATOFMS was operating for 68% and



1 97% of the time at the UB and at the RS site, respectively. It is also worth noting that
2 an inter comparison of the two instruments was attempted at the same site, but this
3 could not be completed due to complex logistical and technical factors.

4 The ATOFMS datasets were imported individually into YAADA (Yet Another
5 ATOFMS Data Analyzer) and single-particle mass spectra were grouped with
6 adaptive resonance theory neural network, ART-2a (Song et al. 1999). The
7 parameters used for ART-2a in this study were: learning rate 0.05, vigilance factor
8 0.85, and 20 iterations. Further details of the parameters can be found elsewhere
9 (Dall'Osto and Harrison 2006; Rebertier and Prather 2007). An ART-2a area matrix
10 (AM) of a particle cluster represents the average intensity for each m/z for all
11 particles within a group. An ART-2a AM therefore reflects the typical mass spectra of
12 the particles within a group.

13

14 **3. Results**

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16 **3.1 ATOFMS particle detection efficiency**

17 Overall, 890,873 particle mass spectra were apportioned at the RS and 221,139 at
18 the UB. This large difference in detected particle numbers is likely a result of the
19 combined effects of the location and detection efficiencies of both instruments. As
20 shown in Figure 1, the number and size distribution of the particles detected by the
21 two mass spectrometers is quite different and reflects their expected performance
22 characteristics (Gard et al., 1997; Su et al., 2004). The instrument with the
23 aerodynamic lens detected considerably more particles below 1 μm , while particles
24 larger than ca. 1.9 μm were only detected with the converging nozzle inlet. By



1 running ART-2a, more than 300 clusters were found initially in both UB and RS
2 datasets. Many were merged if they presented similar temporal trends, size
3 distributions and mass spectra (Dall'Osto and Harrison, 2006). By merging similar
4 clusters, the total number of particle types describing the whole dataset was reduced
5 to 18 and 14 at the RS and UB sites, respectively. Despite the different inlet
6 configurations and sampling locations, the majority of the particles detected at both
7 sites could be described by 10 common ATOFMS particle types, listed in Table 1. A
8 number of other particle types were found only at one of the monitoring sites (8 at the
9 RS site and 4 at the UB site), likely due to the different urban environments, as well
10 as the different detection efficiency of the two instruments (Table 1).

11

12 **3.2 ATOFMS particle types observed at both sites**

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14 **3.2.1 Elemental Carbon (EC)**

15

16 Two main EC particle types, representing together more than 50% of detected
17 particles (58.9% at the UB, 53.8% at the RS), were identified at both sites. For
18 reasons outlined below, they are named EC_Aged_R (Regional) and EC_Aged_L
19 (Local). Both EC particle types presented a fine aerosol size distribution mode at
20 both sites (about 300-500 nm, see Fig. S11). Fig. 2a shows the positive mass
21 spectrum of particle type EC_Aged_R. It is dominated by EC peaks at m/z 12 $[C]^+$, 36
22 $[C_3]^+$, 48 $[C_4]^+$ and 60 $[C_5]^+$. Cluster EC_Aged_L (see Fig. 2b) also shows a similar
23 EC positive mass spectrum pattern, although a strong signal at m/z 39 dominates the
24 positive mass spectrum. This peak is often associated with potassium $[K]^+$, although



1 there may also be a contribution from the organic ion $[C_3H_3]^+$ (Dall’Osto et al., 2009).
2 The EC signals present in the positive mass spectra, and the near-total absence of
3 peaks in the negative mass spectra suggests these EC particles are not freshly
4 emitted. Indeed, Giorio et al. (2012) reported for a regional background site that
5 strong EC peaks in the negative ion mass spectrum are more indicative of fresh
6 emissions while strong EC peaks in the positive mode represent aged elemental
7 carbon. This is consistent with observations which indicate that particle composition
8 affects the ionization and fragmentation pattern of EC (Reinard and Johnson, 2008).
9 The two EC-rich particle types were mainly detected in stagnant air masses. As
10 shown in Fig. SI 1a, an enhancement of these particle types was observed during
11 sulfate-rich air masses (7-10 October 2010) and nitrate-rich stagnant regional air
12 masses (13-17 October 2010), as described in Dall’Osto et al., (2012). The diurnal
13 profiles of these two aged EC particle types (Fig. 3a) can also help to classify their
14 origin. Whilst EC_Aged_R shows little diurnal variation, particle type EC_Aged_L
15 shows a gradual increase during the day, peaks in the afternoon (15:00) and drops to
16 a minimum during night time. The absence of sharp morning peaks indicates that
17 these EC particle types are not associated with primary emission from local traffic.
18 Similarly, the absence of a sharp evening peak also excludes the association with
19 biomass burning, which is not expected at this time of the year and is generally not
20 important in Barcelona (Dall’Osto et al., 2012). Whilst the flat diurnal trend of
21 EC_Aged_R is typical of regional aerosol, particle type EC_Aged_L suggests a local
22 influence, or an enhancement during the warmer part of the day. This assignment is
23 also supported by the observed correlation of the EC types with some secondary
24 inorganic particle types associated with local and regional aerosol sources, described
25 in the next section. Similar conclusions were recently reported by Decesari et al.



1 (2014), where most of the aged EC particles detected with the ATOFMS were related
2 to aged anthropogenic aerosols accumulating in the lower layers of the Po Valley
3 (Italy) overnight.

4 **3.2.2 Secondary nitrate and sulfate**

5 The application of the ART-2a neural network algorithm to the ATOFMS data
6 apportioned two main distinct nitrate particle types, already previously reported
7 (Dall'Osto et al., 2009; Harrison et al., 2012; Decesari et al., 2014). The first ("Loc-
8 NIT" Local-Nitrate; 4.2-7.3% particles by number) appears to be locally produced in
9 urban locations during night time, whilst the second ("Long Range Transport" nitrate,
10 LRT-NIT 4.2-7.3% of particles by number) is regionally transported within the Iberian
11 Peninsula and the rest of Europe. Briefly, particle type Loc-NIT is characteristic of
12 nitrate aerosol in small particles (D_a at about 300–500 nm; Fig. SI2). The average
13 mass spectrum (Fig. 2c) shows a peak at m/z 39 that can be due to potassium,
14 although previous studies (Dall'Osto et al., 2009) suggested that an organic
15 contribution may be also present depending on the m/z 39/41 ratio. The ATOFMS
16 nitrate particle type appears to be associated with local formation processes and
17 occurred in the main at times outside of the long-range transport episode (Fig SI 2b).
18 By contrast, a second nitrate particle type (Long Range Transport nitrate, LRT-NIT) is
19 regionally transported. The average mass spectrum (Fig. 2d) shows that nitrate (m/z
20 -62) is internally mixed with sulfate (m/z -97), ammonium (m/z 18) and both
21 elemental (m/z 36, 48, 60) and organic carbon (m/z 37, 39, 43). As explained in
22 Decesari et al. (2014) these aerosols likely originate from the night-time
23 condensation of nitric acid on EC-containing primary particles (e.g. Shiraiwa et al.,
24 2007). The LRT-NIT particle type is volatile, and partially evaporates during the day



1 leaving a core of about 300 nm mainly composed also of sulfate, elemental and
2 organic carbon (Dall'Osto et al., 2009). Indeed, this can be seen in the LRT-SUL
3 (Long Range Transport - Sulfate) particle type, whose average mass spectrum for
4 this (Fig. 2e) shows peaks due to nitrate (m/z -62), sulfate (m/z -97), elemental
5 carbon (m/z -36, 48, 60) and organic carbon (m/z -39, 43). The LRT-SUL and LRT-
6 NIT diurnal trends are clearly anti-correlated (Fig. 3b), with LRT-SUL concentrations
7 peaking in the afternoon hours. Such behaviour is attributed to the effect of the diel
8 cycle of nitric acid and ammonia condensation/evaporation on the same particle type:
9 during night-time this regional particle type is seen with nitrate, which evaporates
10 during daytime leaving a smaller aerosol core composed of EC and sulfate (Decesari
11 et al., 2014). However, this study shows a novel aspect not found in the London
12 local-regional nitrate study (Dall'Osto et al., 2009) or in the Po Valley (Decesari et al.,
13 2014). During the period 12-14 October, under rainy conditions and with air masses
14 arriving from Europe (Dall'Osto et al., 2012) the transport of LRT-SUL was detected
15 without the nitrate component (LRT-NIT, Fig. S1b). This could be due to a different
16 source of LRT-SUL not linked with LRT-NIT, and previously unobserved.

17

18 **3.2.3 Potassium Organo-Nitrogen (K-CN) particles**

19

20 The K-CN particle type was a minor one, representing only 1.3-2.4% of the total
21 particles analyzed. Fig. 2f shows the average mass spectrum, which features a
22 strong peak at m/z 39 $[K]^+$ in the positive mode, as well as peaks at m/z 113 $[K_2Cl]^+$
23 and m/z -35 $[Cl]^-$ suggesting a biomass burning source (Pastor et al., 2003;
24 Dall'Osto and Harrison, 2006). K-rich particles similar to K-CN have previously been



1 attributed to biomass burning (Silva et al., 1999; Guazzotti et al., 2003) and were
2 found to correlate with gas-phase measurements of acetonitrile, a good biomass-
3 burning tracer. The negative ion mass spectrum shows strong peaks at m/z - 26
4 $[\text{CN}]^-$ and m/z -42 $[\text{CNO}]^-$, indicating that the potassium and chloride is internally
5 mixed with organo-nitrogen species. The size distribution, centred at about 350 nm
6 (Fig. SI 2) also points to a combustion source. Overall, this particle type represented
7 only about 2% of the total particles sampled (Table 1) and its temporal variation is
8 presented in Fig. SI 2d. Its diurnal trend (Fig. 3) tracks the anthropogenic activities of
9 the city of Barcelona, suggesting general minor urban combustion processes. It
10 should be noted that biogenic plant debris also has a similar single particle mass
11 spectrum, with strong signals at m/z 39, - 26 and -42 (Silva et al., 2000). However,
12 vegetative dust is usually internally mixed with sodium and phosphate, and presents
13 aerodynamic diameters above 1 μm (see section 3.3).

14

15 **3.2.4 Fresh and aged sea salt particles**

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17 Two sea salt particle types (fresh and aged) were detected at both sites, accounting
18 for 9.4% and 18.3% of the total particles sampled at RS and UB, respectively (Table
19 1). The average mass spectrum (Fig. 2g) for the particle type assigned to fresh sea
20 salt (labelled NaCl) shows peaks typical of sodium chloride clusters ($[\text{Na}]^+$ (m/z 23),
21 $[\text{K}]^+$ (m/z 39), $[\text{Na}_2]^+$ (m/z 46), $[\text{Na}_2\text{Cl}]^+$ (m/z 81 and 83), $[\text{NaCl}_2]^-$ (m/z 93, 95 and 97),
22 whilst aged sea salt (NaCl-NIT, Fig. 2h) also exhibited nitrate peaks (m/z -46 and
23 m/z -62) reflecting the reaction between NaCl and HNO_3 and the replacement of
24 chloride by nitrate (Gard et al., 1998). Both particle types had mean size distributions



1 above 1 μm (Fig. SI 2 c). NaCl was mainly detected during air masses that had
2 travelled over Mediterranean regions (7-10 and 11-13 October), as shown in Fig SI
3 2e. Interestingly, the diurnal variation of the two NaCl particle types is quite different
4 (Fig. 3). Whilst aged sea salt (NaCl-NIT) does not show a clear trend, the fresh sea
5 salt (NaCl) shows an enhancement in the afternoon, associated with the sea breeze,
6 peaking at 15:00 (Dall'Osto et al., 2012).

7

8 **3.2.5 Iron and Lead containing particles**

9

10 Two common ATOFMS particle types have average mass spectra that were
11 dominated by metals. The first was rich in iron (type Fe, 1.4% of the total particles)
12 and has a spectrum (Fig. 2i) characterised by a strong signal at m/z 56 and weaker
13 features at m/z 73 (iron oxide, $[\text{FeOH}]^+$) and m/z 54 (isotope ^{54}Fe). The negative
14 spectra spectrum has strong features at m/z -46 and -62, indicating that the iron is
15 internally mixed with nitrate. During SAPUSS, the Fe particle type was found to
16 correlate ($R^2=0.75$) with LRT-NIT, which is associated with long range transport of
17 pollutants. The small mode of the Fe particle type (Fig SI 2) reflects the fact that only
18 fine particles were likely to travel long distances relative to the coarser ones which
19 were lost during transport. This is in line with previous ATOFMS field studies showing
20 the transport of iron-containing particles internally mixed with nitrate from continental
21 Europe (Harrison et al., 2012). The flat diurnal profile for the Fe particle type (Fig. 3)
22 also suggests a regional origin.

23 The average mass spectrum of the second metal-rich particle type is shown in Fig.
24 2j. This particle type is labelled Pb since lead is one of the largest contributors in the



1 positive mode, occurring at m/z +206, +207 and +208. Other peaks in the positive ion
2 spectrum include m/z 23 $[\text{Na}]^+$, m/z 56 $[\text{Fe}]^+$, m/z 39 and 113 ($[\text{K}]^+$ and $[\text{K}_2\text{Cl}]^+$). In
3 addition to nitrate (m/z -46 $[\text{NO}_2]^-$ and m/z -62 $[\text{NO}_3]^-$), chloride (m/z -35 ($[\text{Cl}]^-$) was
4 one of the most abundant species in the negative ion mode. A similar particle type
5 containing lead and chloride (Pb-Cl) was detected by ATOFMS in Mexico City and
6 attributed to a waste incinerator source (Moffet et al., 2008), although other studies in
7 the same area attributed lead-containing particles to multiple sources, including
8 trash burning (Salcedo et al., 2010; Hodzic et al., 2012). This ATOFMS particle type
9 was found highly correlated with hourly elemental concentrations determined by
10 PIXE analysis (Dall'Osto et al., 2013), showing that this source of Pb-Cl is a major
11 (82%) source of fine Cl in the urban agglomerate of Barcelona. It is worth noting that
12 whilst the regional particle type Fe was mainly distributed in the fine mode (about
13 300-500 nm), a much larger mode (about 700-900 nm) was observed for this local
14 ATOFMS Pb particle type (see Fig. SI 2 d). Finally, the temporal trend (Fig. S1d) and
15 the diurnal profile (Fig. 3) of the Pb particle type also suggests a local origin, likely
16 related to emissions from urban combustion.

17

18 **3.3 particle types observed at UB**

19 The ATOFMS fitted with converging nozzle inlet located at the UB site detected four
20 particle types that were not observed at the RS site. Each of the particle types make
21 a minor contribution (0.9-1.7%) and overall they represented less than 5% of the total
22 particles sampled at the UB site.

23

24



1 **3.3.1 Vanadium-containing particles (V)**

2 A particle type containing vanadium (m/z 51 $[V]^+$ and m/z 67 $[VO]^+$) Na (m/z 23) and
3 Fe (m/z 56), along with minor peaks due to EC at m/z 36, 48 and 60 was observed at
4 the UB site. A negative ion mass spectrum was not acquired for this particle type,
5 Figure 4a. It has been shown that the condensation of secondary material on soot
6 particles (Moffet and Prather, 2009) and the consequent change in aerosol
7 hygroscopicity (Spencer et al., 2007) can suppress the formation of negative ions in
8 real-time laser desorption/ionisation mass spectrometry, causing many aged EC-
9 containing particles to lack negative mass spectra (Neubauer et al., 1998). The V
10 particle type accounted for only 1% of the particles characterised at the UB, with a
11 unimodal distribution peaking at about 350 nm pointing to a fresh combustion
12 emission source. Field observations have confirmed that ships produce significant
13 amounts of soot, vanadium, nickel, and sulfate (Pattanaik et al., 2007). Previous
14 ATOFMS field measurements also reported that single particles containing organic
15 carbon, vanadium, and sulfate (OC-V-sulfate) resulted from residual fuel combustion
16 (i.e. bunker fuel), whereas high quantities of fresh soot particles represented distinct
17 markers for plumes from distillate fuel combustion (i.e. diesel fuel) from ships as well
18 as trucks in the port area (Healy et al., 2009; Ault et al., 2010). Indeed, the V particle
19 type presented a temporal variation showing a maximum on 3 October 2010, when
20 strong wind was blowing from the port of Barcelona (North African tropical air masses
21 from West, NAF_W, Fig. SI 1 e). This Vanadium particle type is therefore most likely
22 associated with fuel oil and ship emissions from the port. However, it was not
23 detected at the RS site, possibly due to local street canyon effects which restrict
24 transport and mixing in the area (Solazzo et al., 2008).

25



1 **3.3.2 Soil rich in Ca (Soil-Ca)**

2

3 The average ATOFMS mass spectrum of a dust particle type rich in calcium is shown
4 in Fig. 4b, where peaks for calcium (m/z 40 $[Ca]^+$, 56 $[CaO]^+$, 57 $[CaOH]^+$ and 96
5 $[Ca_2O]^+$) are present. Peaks associated with sodium (m/z 23), magnesium (m/z 24,
6 25) and potassium (m/z 39) can also be seen. The Soil-Ca particle class represented
7 0.9% of the particles sampled at the UB site and was found to moderately correlate
8 ($R^2=0.6$) with the urban Dust (Ca) factor obtained by PIXE-PMF analysis (Dall'Osto
9 et al., 2013). The large aerosol size mode above 1 μm (Fig. SI 1e) also points a dust
10 origin. Enhanced levels of the Ca ATOFMS particle type were observed during
11 Atlantic (ATL) air masses (Dall'Osto et al., 2012), likely due to extra resuspension of
12 dust caused by the high wind speeds during these periods. Urban areas in the South
13 of Europe are known to have high dust loadings and three different types of dust
14 were previously reported in the SAPUSS PIXE study (Dall'Osto et al., 2013),
15 representing ca. 25% of the $PM_{2.5}$ mass concentrations measured. It should be noted
16 that this particle type did not possess a negative ion mass spectrum and was also not
17 detected at the RS site, likely for the same reasons explained above (section 3.3.1).

18

19

20 **3.3.3 Saharan dust particles (Saharan-dust)**

21

22 A rarely observed particle type with peaks due to titanium at m/z 48 and m/z 64 ($[Ti]^+$
23 and $[TiO]^+$, respectively) was found to represent 0.9% of the particles sampled at the
24 UB site (Fig. 4c). Additional peaks are associated with other metals including Na (m/z



1 23), Al (m/z 27), Ca (m/z 40, 56, 96) and Fe (m/z 56), as well as minor peaks of
2 silicate at m/z - 60 $[\text{SiO}_2]^-$ and - 76 $[\text{SiO}_3]^-$. These particles presented an aerosol size
3 coarse mode (1-3 μm ; Fig. S12) and were detected mainly during the period 8-10
4 October, when Barcelona experienced air masses originating in the North African
5 Saharan region (Dall'Osto et al., 2012). This particle type was likely not detected at
6 the RS site because the ATOFMS fitted with the aerodynamic focussing lens is not
7 efficient at detecting such coarse particles. With regard to secondary species, it is
8 interesting to note that the dust was internally mixed only with nitrate and not with
9 sulfate (Fig. 4c). Previous studies (Dall'Osto et al., 2010) showed that Saharan dust
10 particles collected near the Cape Verde Islands contained internally mixed nitrate but
11 no sulfate, whilst Saharan dust particles collected on the coast of Ireland showed a
12 very high degree of internally mixed secondary species including nitrate, sulfate and
13 methanesulfonate (Dall'Osto et al., 2004).

14

15 **3.3.4 Vegetative debris (Veg-KP)**

16

17 The ATOFMS has already proven to be a good tool for identifying and separating
18 dust (mainly Ca-rich or Al-Si rich) and biological particles (Ferguson et al. 2004). A
19 particle type dominated by K (m/z 39) and phosphate (m/z -63 and m/z - 79) in the
20 average mass spectrum (Fig. 4d) accounted for 1.7% of the total particles sampled at
21 the UB site. Strong signals at m/z 26 and m/z 42, due to $[\text{CN}]^-$ and $[\text{CNO}]^-$ are also
22 present, as well as peaks due to Na (m/z 23) and Fe (m/z 56) . A coarse aerosol size
23 mode of about 3 μm (Fig. S12) is likely again the reason why this particle type was
24 not detected by the ATOFMS at the RS site. The size and chemical composition



1 suggest a source such as vegetative debris and the particle type is thus labelled Veg
2 K-P. A previous ATOFMS study reported a very similar particle type from samples of
3 leaves collected from a roadside (Schofield, 2004). It is interesting to note that an
4 increase of this vegetative debris particle type was detected under NAF_W air
5 masses in concomitance with Saharan dust particles. Previous atmospheric
6 measurements have shown that the concentration of bacteria over the sea may be
7 much lower than over land, but that higher concentrations of aerosolized
8 microorganisms are generated during dust events compared to clean background
9 marine conditions (Kellog and Griffin 2006, Prospero et al., 2006).

10

11 **3.4 particle types observed at RS**

12

13 The ATOFMS fitted with an aerodynamic focussing lens located at the RS site
14 detected 8 particle types that were not observed at the UB site.

15

16 **3.4.1 Amines**

17 Amines are ubiquitous in the atmospheric environment, and have been detected in
18 marine, urban and rural atmosphere in the gas and particle phases as well as
19 aqueous fog and rain water (Ge et al., 2011). The ATOFMS is a particularly good
20 aerosol instrument for studying amines because the LDI laser wavelength (266 nm)
21 ionizes them very efficiently (Angelino et al., 2001; Healy et al., 2015). During the
22 SAPUSS measurement campaign, four amine particle types were detected at the
23 RS. To the best of our knowledge, this is the first time such a variety of organo-



1 nitrogen particle types has been detected at the single particle level in real time in
2 urban air.

3

4 (a) *Amine (Traf. 58)*: Fig. 5a shows the ATOFMS mass spectra of a particle type
5 called Amine-58; which accounted for 0.8% of the total particles classified at RS. The
6 strongest peak in the positive mass spectra at m/z 58 is likely due to $[\text{C}_2\text{H}_5\text{NH}=\text{CH}_2]^+$.
7 In general, the most important primary fragmentation process occurring for aliphatic
8 amines involves the removal of one of the electrons from the lone pair on N and
9 cleavage of the C-C bond to the nitrogen, with loss of the heavier alkyl group
10 favoured. This cleavage explains the presence of the fragments at m/z 58
11 $[\text{C}_2\text{H}_5\text{NH}=\text{CH}_2]^+$, 72 $[\text{C}_3\text{H}_7\text{NH}=\text{CH}_2]^+$, 86 $[(\text{C}_2\text{H}_5)_2\text{N}=\text{CH}_2]^+$, and 114 $[(\text{C}_3\text{H}_7)_2\text{N}=\text{CH}_2]^+$.
12 The minor peak at m/z 86 was found to correlate ($R^2 = 0.8$) with m/z 58. M/z 86 was
13 suggested as amine species associated with fresh mobile emissions (Angelino et al.,
14 2001). The negative mass spectrum shows nitrate and sulfate internally mixed with
15 this particle type, although the former has a stronger signal than the latter. A size
16 distribution peaking at about 300-400 nm was associated with this particle type. The
17 diurnal variation of Amine (Traf. 58) (Fig. 6a) shows a peak about one hour later than
18 morning and evening rush hour traffic and thus points to relatively fresh vehicular
19 source. In conclusion, both the mass spectra (similar to previous traffic emissions,
20 Angelino et al., 2001) and the temporal trend of this particle type point to traffic
21 emission as a main source.

22

23 (b) *Amine (ETS. 84)* Fig. 5b shows the average mass spectrum of a particle type
24 associated with Environment Tobacco Smoke (ETS) called Amine-84, which



1 represented 0.5% of the total particles classified at RS. In the positive mass
2 spectrum, the peak at m/z 161 is assigned to nicotine, with the main fragment being
3 at m/z 84 (methyl-pyrrole fragment). The negative mass spectrum shows nitrate and
4 sulfate internally mixed with this particle type. The Amine-84 mass spectrum is very
5 similar to that previously reported in Athens ($R^2 = 0.9$; Dall'Osto et al., 2007) and also
6 attributed to tobacco smoke. This particle type generally followed human activity
7 within the city (Fig. 6a), with a main peak in the morning (10 am) and a broader one
8 during the evening (6pm -10 pm). It should be noted that an afternoon peak -
9 potentially related to lunch activities - was not noticed. Spain ranks among the
10 countries with the highest levels of cigarette consumption in the European Union and
11 in the World (WHO, 2010; Ailer et al.; 2013). High outdoor gas-phase nicotine
12 concentrations (0.5 and $1.5 \mu\text{g m}^{-3}$) have been previously measured in Barcelona in
13 summer 2010 (Sureda et al., 2012).

14

15 (c) *Amine (SOA 59)*: Figure 5c shows the average mass spectrum of a particle type
16 called amine-59, which accounted for 0.2% of the total particles classified at RS.
17 Based on previous studies, the strong peak at m/z 59 ($[\text{N}(\text{CH}_3)_3]^+$) is attributed to
18 Trimethylamine (TMA) (Angelino et al., 2001; Zhuang et al., 2012; Healy et al., 2015).
19 Amine (SOA 59) concentration was found particularly enhanced under regional air
20 masses (13-17 October 2010, Fig. SI 1 h) and exhibited a diurnal trend (Fig. 6a)
21 showing maximum concentrations at 3 pm and 10 pm. TMA plays an important role
22 in atmospheric chemistry, yet its pathway towards aerosol is not clear. Rehbein et al.
23 (2011) demonstrated that cloud/fog processing could enhance gas-to-particle
24 partitioning of TMA. TMA can also participate in the formation of secondary organic



1 aerosol. Several studies have shown that gas-phase TMA could form non-salt
2 organic aerosol products through reaction with oxidizing agents (Murphy et al., 2007).

3

4 (d) *Amine (SOA 114)*: Fig. 5d shows the average mass spectrum for a particle type
5 called amine (SOA 114). The main peak is at m/z 114, which may be assigned to
6 dipropylamine or tripropylamine (Angelino et al., 2001; Healy et al., 2015). Minor
7 peaks can also be seen at m/z 58, 74 and 128, which were previously attributed to
8 alkyl ammonium nitrate salt particles formed by reaction of nitric acid and amines
9 (Angelino et al., 2001). The negative ion mass spectrum for Amine (SOA 114) shows
10 a strong peak at m/z -62, confirming the presence of nitrate. Fig. 6a shows that the
11 diurnal profile is similar to that reported for Amine (SOA 59), although with a much
12 stronger peak at 3 pm, during the warmest part of the day for the measurement period
13 (Dall'Osto et al., 2013). This is in line with recent measurements of gaseous amines
14 by ambient pressure proton transfer mass spectrometry (Hanson et al., 2011) which
15 also showed an enhancement of TEA (triethylamine) during the afternoon. Amine
16 (SOA 59) and Amine (SOA 114) presented similar diurnal profiles but their temporal
17 occurrence was not similar. Whilst Amine (SOA 59) was detected mainly during the
18 period of nitrate-rich regional air masses (13-17 October), Amine (SOA 114) was
19 detected mainly under NAF_E air masses rich in sulphate (7-10 October 2010). This
20 observation is discussed further in section 4.

21

22 Overall, the four amine containing particle types were found distributed mainly in the
23 sub micron mode. Amine (Traf. 58) and Amine (ETS 84), which are attributed to
24 primary emissions from traffic and ETS respectively, have similar size distributions



1 which peak around 400-500 nm (Fig. SI 2). Amines (SOA 59) and Amine (SOA 114)
2 were also found in the same size range, but presented a broader shape, suggesting
3 partial condensation of SOA material on existing particles. Previous work of Angelino
4 et al. (2001) used the peak at m/z 118 to track a marker for the oxidation products of
5 alkylamines observed in the condensed phase of aerosol particles, specifically
6 attributed to the oxidation of triethylamine (m/z 118; $[(C_2H_5)_2NCH_2]^+$). The analysis of
7 the ATOFMS dataset for the RS site shows that the majority (about 80%) of amine
8 particle types containing a peak at m/z 118 was found in the tobacco related
9 ATOFMS Amine (ETS 84) class, whereas the remaining 20% was associated with
10 Amine (Traf. 58). This suggests - at least in our study - that m/z 118 cannot be used
11 ly as marker amines produced by secondary processes. Other possible identification
12 of higher mass ions detected and associated with organic nitrate include m/z 74
13 $[(C_2H_4NO_2]^+$, 88 $[(C_3H_6NO_2]^+$ or $[(C_2H_5)_2NO]^+$, 104 $[(C_3H_6NO_3]^+$, and
14 191 $[(C_6H_{11}N_2O_5]^+$. Overall, m/z 104 was found to be to Amine (Traf. 58), whereas
15 m/z 74 and m/z 191 to Amine (SOA 114). Finally, m/z 88 was found associated with
16 both Amine (Traf. 58) and Amine (SOA 114). In summary, organic nitrate peaks were
17 associated with both primary and secondary amine particle types.

18

19 **3.4.1 Organic Carbon (OC) particle types**

20

21 The high efficiency of the ATOFMS equipped with the aerodynamic lens and
22 deployed at the RS allowed us to characterise four particle types rich in organic
23 compounds.

24



1 (a) *Org (Lub. Oil)*: Fig. 7a shows the average mass spectrum of a particle type
2 named "Lub. Oil" (1.8%). The positive ion mass spectrum shows strong peaks for
3 $[Ca]^+$ (m/z 40), and minor ones for OC (m/z 27, 43, 57) and EC (m/z 12, 24, 36, 48,
4 60). The negative mass spectra shows presence of nitrate (m/z -46, -62) and EC (-
5 36, -48, -60). This particle type was only detected in the lower size range of the
6 ATOFMS (ca. 150-400 nm), was found to spike during traffic rush hour times (Fig.
7 6b) and is attributed to lubricating oil from traffic emissions (Spencer et al., 2006;
8 Drewnick et al., 2008).

9

10 (b) *Org. (Aro-NIT)*: A particle type rich in aromatic compounds and internally mixed
11 with nitrate (*Org. (Aro-NIT)*) was found to account for 1.7% of the particles classified
12 at the RS. The average mass spectrum shown in Fig. 7b features strong signals due
13 to organic fragments at m/z 27, 29, 39, 43, 51, 57 and 63. M/z 51 $[C_4H_3]^+$, 63 $[C_5H_3]^+$,
14 77 $[C_6H_5]^+$ and 91 $[C_7H_7]^+$ are indicative of a strong aromatic signature (McLafferty,
15 1993). Common peaks due to nitrate (i.e. m/z -46 $[NO_2]^-$, m/z -62 $[NO_3]^-$) are found in
16 the negative mass spectrum. Fig. 6b shows a daily maximum of this particle type at
17 about 7-8 pm. Two previous ATOFMS studies in Athens (Dall'Osto and Harrison,
18 2006) and London (Dall'Osto and Harrison, 2012) reported very similar particle types
19 (called C-SEC_2 and Na-EC-OC, respectively) which were attributed to condensation
20 of organic compounds onto the particle phase as air temperatures dropped. The
21 secondary particles showed clear internal mixing of organic and inorganic
22 constituents. The size distribution of the Aro-NIT particle type shows a broad mode at
23 about 500-700 nm (Fig. SI 2), also suggesting condensation of species on existing
24 aerosols.



1

2 (c) Org. (*OC-NIT*): A third organic-rich particle type, called OC-NIT (1.5% of the
3 particles classified), was also detected at the RS site. Fig. 7c shows the average
4 positive ion mass spectrum, which contains a strong peak at m/z 39 (K and/or C_3H_3)
5 and minor peaks due to OC and EC, as well as Na (m/z 23). Perhaps the most
6 interesting part of the negative spectrum for this particle type is the presence of
7 signals due to organic nitrogen species (m/z -26 $[CN]^-$ and m/z -42 $[CNO]^-$), not seen
8 in other organic rich particle types shown in Fig. 7. Finally, strong peaks for nitrate
9 (m/z -46, -62) dominate the negative mass spectra. Fig. 6b shows a daily maximum
10 concentration of OC-NIT at 3 pm. The presence of an organo nitrogen component in
11 this particle type, as well as its maximum at 3 pm, is similar to that for the Amine
12 (SOA 59) particle type described in the previous section.

13

14 (d) Org. (*OC-CHO*): This particle type presents a positive mass spectrum (Figure
15 7d), with strong peaks at m/z 27 $[C_2H_3]^+$ and m/z 43 $[(CH_3)CO]^+$ usually associated
16 with oxidized secondary organic aerosol (Sullivan and Prather, 2005). Peaks at m/z
17 29 $[C_2H_5]^+$ and m/z 41 $[C_3H_5]^+$ confirm the strong hydrocarbon-like nature of the
18 particles (Spencer et al., 2006), while weak peaks at m/z 51 $[C_4H_3]^+$, 63 $[C_5H_3]^+$, 77
19 $[C_6H_5]^+$ and 91 $[C_7H_7]^+$ also indicate an aromatic contribution (McLafferty, 1993). A
20 negative ion mass spectrum was often not recorded for this particle type. However,
21 some of the particles did generate one and an example is shown in Fig. SI 3 d. The
22 presence of peaks at m/z -45, -59 and -71 are likely to be due to the formate
23 $[CHO_2]^-$, acetate $[C_2H_3O_2]^-$ and propionate $[C_3H_5O_2]^-$ ions, respectively. Figure 6b
24 shows that in contrast to the other three OC-rich particle types, OC-CHO shows a



1 complex diurnal profile. The number of counts per hour increases from 6-8 am and
2 then remains stable for about 5 hours. A peak is observed around 3 pm, the hottest
3 part of the day, followed by a second peak at about 10-11pm. This diurnal profile is
4 similar to that of Amine (SOA 114) , which is also of secondary origin.

5 Overall these four OC-rich particle types indicate that a number of processes
6 and sources are likely producing oxidised organic aerosols. There are at least four
7 main peaks during the day: a morning traffic rush hour (9-10 am), an afternoon one
8 during the hottest part of the day (3 pm), and two evening ones at 8 pm (sun set) and
9 at about 10-11pm.

10

11 **4. Discussion and conclusions**

12

13 During the SAPUSS intensive field study two ATOFMS instruments were deployed
14 simultaneously. The ATOFMS deployed at the RS site was equipped with an
15 aerodynamic lens inlet system, allowing characterisation of primary traffic aerosols as
16 well as other primary and secondary aerosols affecting this heavily urbanised area of
17 Barcelona. This type of ATOFMS (Su et al., 2004) has a very high efficiency in
18 sampling aerosols (more than 1,000,000 single particle mass spectra were collected
19 at the RS), particularly for sub micron particles in the size range 300-700 nm. The
20 ATOFMS deployed at the UB site was equipped with a converging nozzle inlet
21 system (Gard et al., 1997), which has a low aerosol collection efficiency (Dall'Osto et
22 al., 2006) but it is particularly well suited for sampling coarser aerosols in the size
23 range 800-2000 nm.



1 Overall, ten particles types were detected at both sites (Table 1). Two of these
2 particle types, composed of EC internally mixed with secondary inorganic species,
3 described more than half of the classified single particle mass spectra. EC_aged_R.
4 was found accumulating within stagnant air masses, with a flat diurnal profile and
5 suggesting a certain physico-chemical stability. In contrast, a more local but
6 processed form of EC (EC_aged_L.) was found to possess a finer submicron mode,
7 and enhanced concentrations during afternoon periods.

8 Two different types of nitrate-dominated aerosols were observed, in line with a
9 previous ATOFMS study in London (Dall'Osto et al., 2009). The first (LRT_NIT) was
10 attributed to regional nitrate (accumulation mode, volatile, more NH_4NO_3 / $(\text{NH}_4)_2\text{SO}_4$
11 type) and the second (Loc_NIT) was assigned to local nitrate (ultrafine mode, less
12 volatile, more OC- NO_3 type). An enhancement of local nitrate was found under warm
13 and humid NAF_E air masses, indicating that meteorological conditions (i.e. high RH)
14 or aerosol composition (i.e. sea salt, Saharan dust) could affect local urban nitrate
15 aerosol production.

16 Two types of sea salt particles were also identified at both sites. Fresh sea salt
17 particles (NaCl) showed a peak in the diurnal profile at 3 pm, related to the sea
18 breeze and enhanced under NAF_E air masses. Aged sea salt particles (NaCl-NIT)
19 were mainly observed during more anthropogenically influenced air masses.

20 Two types of particles rich in metallic elements were found at both monitoring sites.
21 One, rich in iron and internally mixed with nitrate, was found to be distributed in the
22 fine accumulation mode at about 400 nm and related to regional air masses. This
23 observation supports previous findings (Dall'Osto et al., 2010, Harrison et al., 2012)
24 that showed fine iron-containing aerosols are able to travel long distances and are



1 thus related to aged air masses. Other studies have reported anthropogenic Fe-
2 containing particles internally mixed with secondary species such as sulfate (Furutani
3 et al., 2011; Moffet et al., 2012) originating from coal combustion in Asian continental
4 outflows. This study shows that - within the European continental outflow - Fe-
5 containing particles are instead internally mixed with nitrate. This observed difference
6 is likely due to emissions from coal combustion in Asia which are rich in SO₂, as
7 opposed to European air masses which are relatively higher in NO₂ and lower in SO₂.
8 Bio-available iron from atmospheric aerosol is an essential nutrient that can control
9 ocean productivity (Baker and Croot, 2010). Hence, it can impact the global carbon
10 budget and climate. There are also large uncertainties in the origin of the aerosol
11 nitrogen matter which may be enhancing ocean productivity (Duce et al. 2008). The
12 fact iron is internally mixed with nitrate points to an urban source (i.e. more traffic
13 than industrial activities), suggesting transported submicron urban particles can be a
14 source of both iron and nitrogen nutrients for the oceans. However, it is worth noting
15 that Cu-rich traffic related aerosols can also have a negative effect on marine
16 phytoplankton over a vast region of the western Mediterranean Sea (Jordi et al.,
17 2012).

18 A second metallic particle type rich in lead and chloride was identified. This particle
19 type was related to more local sources, presenting sharp spikes in concentration. It is
20 interesting to note this particle type was found correlated with hourly elemental mass
21 concentrations determined by PIXE analysis (Dall'Osto et al., 2013) showing this
22 particle type can be a major source of submicron chloride in the urban area of
23 Barcelona.



1 The ATOFMS equipped with the converging nozzle inlet at the UB site detected four
2 further different particle types. Two types of dust particles were found, both occurring
3 mainly in the coarse mode ($>1 \mu\text{m}$): one type, as attributed to Saharan Dust, was
4 characterized by an aluminium/silicon signature, while the other type, with a more
5 local origin (Ca dust), was characterized by a Ca-rich composition. Two other minor
6 coarse particle types were characterised: vegetative debris (Veg-KP, rich in
7 potassium and phosphate), found enhanced during tropical NAF_E air masses; and
8 vanadium-containing particles (Oil-V), and related to shipping/oil combustion
9 activities in the port of Barcelona. Overall, high concentrations of vanadium were only
10 observed on one day (3 October), indicating that shipping emissions was a minor
11 source of aerosol at the UB site.

12 Eight particle types were detected by the ATOFMS with aerodynamic focussing lens
13 at the RS site. Overall, these particle types described less than 10% of the aerosol
14 population, but their mass spectra, as well as their peculiar diurnal profiles, allow us
15 to advance our understanding of the OC-ON-Nitrate mixing state of urban aerosols.

16 Four particle types contain amines, which, in addition to ammonia, are important
17 atmospheric bases (Ge et al., 2011). Urban concentrations of ammonia in Barcelona
18 are higher than those reported in similar urban background sites in Europe,
19 especially in summer (Reche et al., 2012). Conversely, in winter, levels of ammonia
20 were higher at traffic-affected sites, suggesting a contribution from vehicle emissions
21 (Reche et al., 2012). In comparison, the sources, atmospheric transformation and
22 sinks of amines are more poorly characterised. Overall, both primary (Amine Traf. 58
23 and Amine ETS 84) and secondary (Amine SOA 59 and Amine SOA 114) sources of
24 amine-containing particles were identified during SAPUSS. The most abundant



1 amine particle type (Amine Traf. 58, 0.8% of total particles at RS) was attributed to
2 traffic activities and the second most abundant (Amine ETS 84, 0.5% of the total
3 particles at RS) was also associated with environmental tobacco smoke. Concurrent
4 SAPUSS measurements (Alier et al., 2013) of nicotine concentrations were much
5 higher at the RS site (58 ng m^{-3}) than at UB (7 ng m^{-3}), pointing to a significant
6 outdoor cigarette consumption in the city centre. A recent study (Sleiman et al., 2010)
7 found that nicotine can contribute significantly to the formation of urban SOA though
8 reaction with ozone. The results of this study therefore suggest that third-hand
9 cigarette smoke may be a source of nitrogen-containing particles in Barcelona and
10 similar cities in Southern Europe where smoking is prominent.

11 Amines were also related to secondary aerosol production, although a very complex
12 dynamic was found associated with their occurrences. They were found to peak
13 during the warmest part of the day (3 pm) and during evening times (10-11 pm).
14 Amine (SOA 59), was found particularly enhanced in regional air masses (13-17
15 October 2010) when nitrate concentrations were also high (Dall'Osto et al., 2013). By
16 contrast, Amine (SOA114) was more abundant in NAF_E humid air masses (7-11
17 October 2010). This latter type of SOA had been internally mixed with nitrate,
18 suggesting aminium salt formation under such specific conditions. Previous ATOFMS
19 studies reported that most of the amines volatilised during cold seasons, whereas
20 during summer most were in the form of low-volatility aminium nitrate and sulphate
21 salts when particle acidity was higher (Pratt et al., 2009); supporting laboratory
22 studies reporting that aerosol containing non-salt organic amines are more stable
23 than nitrate salts (Murphy et al., 2007). Overall, amines can undergo oxidation by
24 OH, O₃, and/or NO₃ to form amides, nitramines, and imines, which can also partition
25 to the particle phase (Murphy et al., 2007; Silva et al., 2008, Healy et al., 2015).



1 Rehbein et al. (2011) combined field and laboratory work to demonstrate that high
2 relative humidity greatly enhances the gas-to-particle partitioning and subsequent
3 aqueous acid–base reaction of amines. During the SAPUSS study, we find aminium
4 nitrate salts on coarse particles containing sulfate, suggesting that heterogeneous
5 reactions occur during the warmest part of the day. This observation is in contrast to
6 the study of Day et al. (2010), who reported organo-nitrates when sulfate aerosols
7 and humidity were low, although the study site was less urbanised than the one used
8 in this study.

9 Finally, some consideration should be given to the four specific organic particle types
10 detected at the RS. One (Org. (Lub. Oil)) was found to be related to primary
11 lubricating oil traffic emissions. More difficult is the attribution of the remaining three,
12 which each contain an internal mixture of OC and nitrate. This is not surprising, given
13 the fact the urban atmosphere is heavily contaminated by traffic emissions, the main
14 producers of the two chemical species (Dall’Osto et al., 2012a).

15 Two different types of organic carbon / nitrate particle types were found. One (OC-
16 NIT) was found spiking in the afternoon. By contrast, a nitrate with a strong aromatic
17 signature (OC-Aro-NIT) was found mainly during night time (80% of the time), and
18 showing a sharp concentration peak at 7-8 pm. These two ATOFMS particle types
19 add complexity to the local-regional inorganic nitrate containing particle types, and
20 further studies are needed to correctly apportion the local urban nitrate component.

21 The fourth organic particle type (OC-CHO) was found rich in oxidised organic carbon
22 and associated with nitrate from a traffic source. In a previous ATOFMS study,
23 considerable effort was made to apportion cooking-related particle types (Dall’Osto
24 and Harrison, 2012). However, only a particle type exhibiting maximum frequency



1 during the warmest part of the day and associated with secondary aerosol production
2 from traffic-related semi-volatile aromatic compounds was found. During this study,
3 we were again not able to associate a specific particle type with cooking activities. In
4 a companion SAPUSS study, Alier et al. (2013) reported an aerosol source formed
5 mainly by C₇-C₉ dicarboxylic acids and detected especially during daytime, highly
6 dynamic, dependent on air masses and pointing to a secondary organic component
7 driven by primary urban sources including cooking and traffic (mainly gasoline)
8 activities. O'Brien et al. (2013) reported a detailed mass spectrometry
9 characterisation of the urban aerosol in an urban background environment.
10 Compounds containing only carbon, hydrogen, and oxygen (CHO), and nitrogen-
11 containing organic compounds (NOC) were found, showing that both photo-oxidation
12 and ammonia chemistry may play a role in forming the compounds observed in the
13 mixed urban-rural environment. NOC had precursor product pairs consistent with
14 imidization and cyclization reactions, suggesting that part of the aromatic compounds
15 detected during SAPUSS may also be formed by cyclization and not only by
16 condensation of aromatic VOC.

17

18 In summary, the two ATOFMS instruments deployed during the SAPUSS field
19 measurement study showed that the urban atmosphere contains a complex mixture
20 of aerosol particles emitted from a variety of sources and formed via numerous
21 atmospheric processes. We have identified 22 different particle types, characterised
22 by specific single particle mass spectra and temporal trends. Particle composition is
23 strongly affected by the prevailing air masses. The ATOFMS measurements also
24 provided some novel information on the mixing state of organic carbon and nitrate in



1 urban aerosols, again highlighting the complex nature of the roles played by both
2 primary sources and in situ chemical processing in affecting aerosol composition.

3

4

5

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7

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13 C02-01-AERTRANS). The SAPUSS team is also acknowledged.

14

15

16 **References**

17

18 Abbatt J. P. D., Leea, A. K. Y. and Thornton J.A. Quantifying trace gas uptake to
19 tropospheric aerosol: recent advances and remaining challenges. *Chem. Soc. Rev.*,
20 2012, 41, 6555–6581, 2012

21

22 Alier, M., van Drooge, B. L., Dall'Osto, M., Querol, X., Grimalt, J. O., and Tauler, R.:
23 Source apportionment of submicron organic aerosol at an urban background and a
24 road site in Barcelona (Spain) during SAPUSS, *Atmos. Chem. Phys.*, 13, 10353-
25 10371, doi:10.5194/acp-13-10353-2013, 2013.

26

27 Allan, D., Williams, P.I., Morgan, W.T., Martin, C.L., Flynn, M.J., Lee, J., Nemitz, E.,
28 Phillips, G.J., Gallagher, M.W., Coe, H.: Contributions from transport, solid fuel



- 1 burning and cooking to primary organic aerosols in two UK cities, Atmos. Chem.
2 Phys. 10, 2, 647-668, 2010.
3
- 4 Angelino, S.; Suess, D. T.; Prather, K. A. Formation of aerosol particles from
5 reactions of secondary and tertiary alkylamines: Characterization by aerosol time-of-
6 flight mass spectrometry. Environ. Sci. Technol, 35 (15), 3130–3138, 2001.
7
- 8 AQEG (2005). "Particulate matter in the UK." Defra, London.
9
- 10 Ault, A. P., Gaston, C. J., Wang, Y., Dominguez, G., Thiemens, M. H., and Prather,
11 K. A.: Characterization of the Single Particle Mixing State of Individual Ship Plume
12 Events Measured at the Port of Los Angeles, Environ. Sci. Technol., 44, 1954–1961,
13 doi:10.1021/es902985h, 2010.
14
- 15 Baker, A. Croot. P. L. Atmospheric and marine controls on aerosol iron solubility in
16 seawater, Marine Chemistry, 120, 4-13, 2010.
17
18
- 19 Benton, A. K., Langridge, J. M., Ball, S. M., Bloss, W. J., Dall'Osto, M., Nemitz, E.,
20 Harrison, R. M., and Jones, R. L.: Night-time chemistry above London:
21 measurements of NO₃ and N₂O₅ from the BT Tower, Atmos. Chem. Phys., 10,
22 9781-9795, doi:10.5194/acp-10-9781-2010, 2010
23
- 24 Brown, S. S., deGouw, J. A., Warneke, C., Ryerson, T. B., Dube, W. P., Atlas, E.,
25 Weber, R. J., Peltier, R. E., Neuman, J. A., Roberts, J. M., Swanson, A., Flocke, F.,
26 McKeen, S. A., Brioude, J., Sommariva, R., Trainer, M., Fehsenfeld, F. C., and
27 Ravishankara, A. R.: Nocturnal isoprene oxidation over the Northeast United States in
28 summer and its impact on reactive nitrogen partitioning and secondary organic
29 aerosol, Atmos. Chem. Phys., 9, 3027–3042, doi:10.5194/acp-9-3027-2009, 2009.
30
- 31 Canagaratna, M. R., Jayne, J. T., Jimenez, J. L., Allan, J. D., Alfarra, M. R., Zhang,
32 Q., Onasc, T. B., Drewnick, F., Coe, H., Middlebrook, A., Delia, A., Williams, L. R.,
33 Trimborn, A. M., Northway, M. J., DeCarlo, P. F., Kolb, C. E., Davidovits, P. and
34 Worsnop D. R.: Chemical and microphysical characterization of ambient aerosols
35 with the aerodyne aerosol mass spectrometer, Mass Spectrom. Rev., 26, 185–222,
36 2007.
37
- 38 Dall'Osto, M., Beddows, D.C.S., Kinnersley, R.P., Harrison, R.M., Donovan, R.J. and
39 Heal M.R. Characterization of individual airborne particles by using Aerosol Time-of-
40 Flight Mass Spectrometry (ATOFMS) at Mace Head, Ireland. Journal of Geophysical
41 Research 109, D21302, doi: 10.1029/2004/JD004747, 2004.
42
- 43 Dall'Osto, M., Harrison, R. M., Beddows, D. C. S., Freney, E. J., Heal, M. R. and
44 Donovan R. J.: Single-particle detection efficiencies of aerosol time-of-flight mass
45 spectrometry during the North Atlantic marine boundary layer experiment, Environ.
46 Sci. Technol., 40, 5029-5035, 2006.
47
- 48 Dall'Osto, M. and Harrison R. M.: Chemical Characterisation of single airborne
49 particles in Athens (Greece) by ATOFMS, Atmos. Environ., 40, 7614-7631, 2006.



1
2 Dall'Osto, M., R. M. Harrison, E. Charpantidou, G. Loupa and S. Rapsomanikis.
3 "Characterisation of indoor airborne particles by using real-time aerosol mass
4 spectrometry." *Science Of The Total Environment* 384(1-3): 120-133, 2007.
5
6
7 Dall'Osto, M.; R. M. Harrison, E. J. Highwood, C. O'Dowd, D. Ceburnis, X. Querol, E.
8 P. Achterberg, Variation of the mixing state of Saharan dust particles with
9 atmospheric transport, *Atmospheric Environment*, Volume 44, Issue 26, 3135-3146,
10 ISSN 1352-2310, DOI: 10.1016/j.atmosenv.2010.05.030, 2010.
11
12
13 Dall'Osto, M., Harrison, R. M., Coe, H., Williams, P. I., and Allan, J. D.: Real time
14 chemical characterization of local and regional nitrate aerosols, *Atmos. Chem. Phys.*,
15 9, 3709-3720, 2009.
16
17 Dall'Osto, M., Querol, X., Alastuey, A., Minguillon, M. C., Alier, M., Amato, F., Brines,
18 M., Cusak, M., Grimalt, J. O., Karanasiou, A., Moreno, T., Pandolfi, M., Pey, J.,
19 Reche, C., Ripoll, A., Tauler, R., Van Drooge, B. L., Viana, M., Harrison, R.M., Gietl,
20 J., Beddows, D., Bloss, W., O'Dowd, C., Ceburnis, D., Martucci, G., Ng, S., Worsnop,
21 D., Wenger, J., Mc Gillcuddy, E., Sudou, J., Healy, R., Lucarelli, F., Nava, S.,
22 Jimenez, J.L., Gomez Moreno, F., Artinano, B., Prevot, A.S., Pfaffenberger, L., Frey,
23 S., Wilsenack, F., Casabona, D., Jimenez-Guerrero, P., Gross, D., Cotz, N.:
24 Presenting SAPUSS: Solving Aerosol Problem by using synergistic strategies at
25 Barcelona, Spain, *Atmos. Chem. Phys. Discuss.*, 12, 18741-18815, 2012a.
26
27 Dall'Osto, M. and Harrison, R. M.: Urban organic aerosols measured by single
28 particle mass spectrometry in the megacity of London, *Atmos. Chem. Phys.*, 12,
29 4127-4142, doi:10.5194/acp-12-4127-2012, 2012b.
30
31 Dall'Osto, M., Querol, X., Amato, F., Karanasiou, A., Lucarelli, F., Nava, S., Calzolari,
32 G., and Chiari, M.: Hourly elemental concentrations in PM_{2.5} aerosols sampled
33 simultaneously at urban background and road site during SAPUSS – diurnal
34 variations and PMF receptor modelling, *Atmos. Chem. Phys.*, 13, 4375-4392,
35 doi:10.5194/acp-13-4375-2013, 2013a.
36
37 Dall'Osto, M., Ovadnevaite, J., Ceburnis, D., Martin, D., Healy, R. M., O'Connor, I. P.,
38 Kourtchev, I., Sodeau, J. R., Wenger, J. C., and O'Dowd, C.: Characterization of
39 urban aerosol in Cork city (Ireland) using aerosol mass spectrometry, *Atmos. Chem.*
40 *Phys.*, 13, 4997-5015, doi:10.5194/acp-13-4997-2013, 2013b.
41
42 Day, D. A., Liu, S., Russell, L. M. and Ziemann, P. J. Organonitrate group
43 concentrations in submicron particles with high nitrate and organic fractions in
44 coastal southern California. *Atmospheric Environment* 44:1970-1979, 2010.



1

2 Decesari, S., Allan, J., Plass-Duelmer, C., Williams, B. J., Paglione, M., Facchini, M.
3 C., O'Dowd, C., Harrison, R. M., Gietl, J. K., Coe, H., Giulianelli, L., Gobbi, G. P.,
4 Lanconelli, C., Carbone, C., Worsnop, D., Lambe, A. T., Ahern, A. T., Moretti, F.,
5 Tagliavini, E., Elste, T., Gilge, S., Zhang, Y., and Dall'Osto, M.: Measurements of the
6 aerosol chemical composition and mixing state in the Po Valley using multiple
7 spectroscopic techniques, *Atmos. Chem. Phys.*, 2014, 14, 12109-12132,
8 doi:10.5194/acp-14-12109-2014.

9

10

11 Drewnick, F., Dall'Osto, M., and Harrison, R. M.: Characterization of aerosol particles
12 from grass mowing by joint deployment of ToF-AMS and ATOFMS instruments,
13 *Atmos. Environ.*, 42, 3006–3017, 2008.

14

15 Duce, R.A., La Roche, J., Altieri, K., Arrigo, K.R., Baker, A.R., Capone, D.G., Cornell,
16 S., Dentener, F., Galloway, J., Ganeshram, R.S., Geider, R.J., Jickells, T., Kuypers,
17 M.M., Langlois, R., Liss, P.S., Liu, S.M., Middleburg, J.J., Moore, C.M., Nickovic, S.,
18 Oschlies, A., Pedersen, T., Prospero, J., Schlitzer, R., Seitzinger, S., Sorensen, L.L.,
19 Uematsu, M., Ulloa, O., Voss, M., Ward, B., Zamora, L.: Impacts of atmospheric
20 anthropogenic nitrogen on the open ocean. *Science* 320, 893-897, 2008

21

22

23 Fergenson, D. P., Pitesky, M. E., Tobias, H. J., Steele, P. T., Czerwieniec, G. A.,
24 Russell, D. H., Lebrilla, C. B., Horn, J. M., Coffee, K. R., Srivastava, A., Pillai, S. P.,
25 Shih, M.-T. P., Hall, H. L., Ramponi, A. J., Chang, J. T., Langlois, R. G., Estacio, P.
26 L., Hadley, R. T., Frank, M. and Gard, E. E., Reagentless Detection and Classification
27 of Individual Bioaerosol Particles in Seconds, *Anal. Chem.*, 76, 373-378, 2004.

28

29

30 Furutani, H., Jung, K., Miura, A., Takami, S., Kato, Y., Kajii, and M. Uematsu,
31 Single - particle chemical characterization and source apportionment of
32 iron - containing atmospheric aerosols in Asian outflow, *J. Geophys. Res.*, 116,
33 D18204, doi:10.1029/2011JD015867, 2011.

34

35 Gard, E., Mayer, J. E., Morrical, B. D., Dienes, T., Fergenson, D. P. and Prather K.
36 A.: Real-time analysis of individual atmospheric aerosol particles: Design and
37 performance of a portable ATOFMS, *Anal. Chem.*, 69, 4083-4091, 1997.

38

39

40 Gard, E. E., M. J. Kleeman, D. S. Gross, L. S. Hughes, J. O. Allen, B. D. Morrical, D.
41 P. Fergenson, T. Dienes, M. E. Galli, R. J. Johnson, G. R. Cass, and K. A. Prather,
42 Direct observation of heterogeneous chemistry in the atmosphere, *Science*, 279,
43 1184–1187, 1998.

44

45

46 Ge, X., Wexler, A. S., and Clegg, S. L.: Atmospheric amines – Part I. A review,
47 *Atmos. Environ.*, 45, 524–546, doi:10.1016/j.atmosenv.2010.10.012, 2011.

48



- 1
2 Giorio, C., Tapparo, M., Dall'Osto, M., Harrison, Roy M., Beddows, D.C.S, Di Marco,
3 C., Nemitz, E. Comparison of three techniques for analysis of data from an Aerosol
4 Time-of-Flight Mass Spectrometer, *Atmospheric Environment*, Volume 61, Pages
5 316-326, 2012
6
7
8 Goldstein, A. H. and Galbally, I. E.: Known and unexplored organic constituents in
9 the earth's atmosphere, *Environ. Sci. Technol.*, 41, 1514-1521, 2007.
10
11 Gross, D. S., Galli, M. E., Silva, P. J., and Prather, K. A. Relative Sensitivity Factors
12 for Alkali Metal and Ammonium Cations in Single Particle Aerosol Time-of-flight Mass
13 Spectra, *Anal. Chem.*, 72, 416–422, 2000.
14
15 Guazzotti, S.A., Suess, D.T., Coffee, K.R., Quinn, P.K., Bates, T.S., Wisthaler, A.,
16 Hansel, A., Ball, W.P., Dickerson, R.R., Neususs, C., Crutzen, P.J., Prather, K.A..
17 Characterization of carbonaceous aerosols outflow from India and Arabia:
18 biomass/biofuel burning and fossil fuel combustion. *Journal of Geophysical*
19 *Research—Atmospheres* 108 (D15) (Article no. 4485), 2003.
20
21 Hanson, D. R., McMurry, P. H., Jiang, J. K., Huey, G., and Tanner, D.: Ambient
22 Pressure Proton Transfer mass spectrometry: detection of ammonia and amines,
23 *Environ. Sci. Technol.*, 45, 8881–8888, 2011.
24
25 Harrison, R. M., Dall'Osto, M., Beddows, D. C. S., Thorpe, A. J., Bloss, W. J., Allan,
26 J. D., Coe, H., Dorsey, J. R., Gallagher, M., Martin, C., Whitehead, J., Williams, P. I.,
27 Jones, R. L., Langridge, J. M., Benton, A. K., Ball, S. M., Langford, B., Hewitt, C. N.,
28 Davison, B., Martin, D., Petersson, K. F., Henshaw, S. J., White, I. R., Shallcross, D.
29 E., Barlow, J. F., Dunbar, T., Davies, F., Nemitz, E., Phillips, G. J., Helfter, C., Di
30 Marco, C. F., and Smith, S.: Atmospheric chemistry and physics in the atmosphere of
31 a developed megacity (London): an overview of the REPARTEE experiment and its
32 conclusions, *Atmos. Chem. Phys.*, 12, 3065-3114, doi:10.5194/acp-12-3065-2012,
33 2012.
34
35
36 Healy, R. M., O'Connor, I. P., Hellebust, S., Allan, A., Sodeau, J. R., and Wenger,
37 J. C.: Characterisation of single particles from in-port ship emissions, *Atmos.*
38 *Environ.*, 43, 6408–6414, 2009.
39
40 Healy, R. M., Sciare, J., Poulain, L., Crippa, M., Wiedensohler, A., Prévôt, A. S. H.,
41 Baltensperger, U., Sarda-Estève, R., McGuire, M. L., Jeong, C.-H., McGillicuddy, E.,
42 O'Connor, I. P., Sodeau, J. R., Evans, G. J., and Wenger, J. C.: Quantitative
43 determination of carbonaceous particle mixing state in Paris using single-particle
44 mass spectrometer and aerosol mass spectrometer measurements, *Atmos. Chem.*
45 *Phys.*, 13, 9479-9496, doi:10.5194/acp-13-9479-2013, 2013
46



1 Healy, R. M., Riemer, N., Wenger, J. C., Murphy, M., West, M., Poulain, L.,
2 Wiedensohler, A., O'Connor, I. P., McGillicuddy, E., Sodeau, J. R., and Evans, G. J.:
3 Single particle diversity and mixing state measurements, *Atmos. Chem. Phys.*, 14,
4 6289-6299, doi:10.5194/acp-14-6289-2014, 2014.

5

6 Healy RM, Evans GJ, Murphy M, Sierau B, Arndt J, McGillicuddy E, O'Connor IP,
7 Sodeau JR, Wenger JC. Single-particle speciation of alkylamines in ambient aerosol
8 at five European sites. *Anal Bioanal Chem.* Aug;407(20):5899-909. doi:
9 10.1007/s00216-014-8092-1, 2015.

10

11 Hodzic A, Wiedinmyer C, Salcedo D, Jimenez JL. Impact of trash burning on air
12 quality in Mexico City. *Environ Sci Technol*, 1;46(9):4950-7, 2012

13

14 Jeong, C.-H., McGuire, M. L., Godri, K. J., Slowik, J. G., Rehbein, P. J. G., and
15 Evans, G. J.: Quantification of aerosol chemical composition using continuous single
16 particle measurements, *Atmos. Chem. Phys.*, 11, 7027-7044, doi:10.5194/acp-11-
17 7027-2011, 2011

18

19 Jimenez, J. L., Canagaratna, M. R., Donahue, N. M., Prevot, A. S. H., Zhang, Q.,
20 Kroll, J. H., DeCarlo, P. F., Allan, J. D., Coe, H., Ng, N. L., Aiken, A. C., Docherty, K.
21 S., Ulbrich, I. M., Grieshop, A. P., Robinson, A. L., Duplissy, J., Smith, J. D., Wilson,
22 K. R., Lanz, V. A., Hueglin, C., Sun, Y. L., Tian, J., Laaksonen, A., Raatikainen, T.,
23 Rautiainen, J., Vaattovaara, P., Ehn, M., Kulmala, M., Tomlinson, J. M., Collins, D.
24 R., Cubison, M.J., Dunlea, E. J., Huffman, J. A., Onasch, T. B., Alfarra, M. R.,
25 Williams, P. I., Bower, K., Kondo, Y., Schneider, J., Drewnick, F., Borrmann, S.,
26 Weimer, S., Demerjian, K., Salcedo, D., Cottrell, L., Griffin, R., Takami, A., Miyoshi,
27 T., Hatakeyama, S., Shimono, A., Sun, J. Y., Zhang, Y. M., Dzepina, K., Kimmel, J.
28 R., Sueper, D., Jayne, J. T., Herndon, S. C., Trimborn, A. M., Williams, L. R., Wood,
29 E. C., Middlebrook, A. M., Kolb, C.E., Baltensperger, U., and Worsnop, D. R.:
30 Evolution of organic aerosols in the atmosphere, *Science*, 326, 1525–1529, 2009.

31

32 Jordi, A., Basterretxea, G., Tovar-Sánchez, A., Alastuey, A., Xavier Querol X. Copper
33 aerosols inhibit phytoplankton growth in the Mediterranean Sea PNAS 2012 ;
34 published ahead of print December 10, 2012, doi:10.1073/pnas.1207567, 2012

35

36 Kellogg, C. A. and Griffin, D.W., *Aerobiology and the global transport of desert dust*,
37 *Trends Ecol Evol*, 21, 638–644, doi: 10.1016/j.tree.2006.07.004, 2006.

38

39

40 Laskin, A., Laskin, J., and Nizkorodov, S. A.: Mass spectrometric approaches for
41 chemical characterization of atmospheric aerosols: critical review of the most recent
42 advances, *Environ. Chem.*, 9, 163–189, 2012

43

44 McLafferty, F. W.: *Interpretation of Mass Spectra*, 3rd Edn., p. 303, 1993.

45

46 Mc Murry P.H. A review of atmospheric aerosol measurementsq *Atmos, Environ*
47 *34:1959–1999*, 2000



1

2 Mohr, C., DeCarlo, P. F., Heringa, M. F., Chirico, R., Slowik, J. G., Richter, R.,
3 Reche, C., Alastuey, A., Querol, X., Seco, R., Penuelas, J., Jimenez, J. L., Crippa,
4 M., Zimmermann, R., Baltensperger, U., and Prevot, A. S. H.: Identification and
5 quantification of organic aerosol from cooking and other sources in Barcelona using
6 aerosol mass spectrometer data, *Atmos. Chem. Phys.*, 12, 1649–1665,
7 doi:10.5194/acp-12-1649-2012, 2012

8

9 Moffet, R. C., de Foy, B., Molina, L. T., Molina, M. J., and Prather, K. A.:
10 Measurement of ambient aerosols in northern Mexico City by single particle mass
11 spectrometry, *Atmos. Chem. Phys.*, 8, 4499–4516, doi:10.5194/acp-8-4499-2008,
12 2008

13

14 Moffet, R. C. and Prather, K. A.: In-situ measurements of the mixing state and optical
15 properties of soot with implications for radiative forcing estimates, *Proc. Natl. Acad.*
16 *Sci. U. S. A.*, 106, 11872–11877, 2009.

17

18 Moffet, R. C., H. Furutani, T. C. Rödel, T. R. Henn, P. O. Sprau, A. Laskin, M.
19 Uematsu, and M. K. Gilles, Iron speciation and mixing in single aerosol particles from
20 the Asian continental outflow, *J. Geophys. Res.*, 117, D07204,
21 doi:10.1029/2011JD016746, 2012

22

23 Murphy, S. M., Sorooshian, A., Kroll, J. H., Ng, N. L., Chhabra, P., Tong, C., Surratt,
24 J. D., Knipping, E., Flagan, R. C., and Seinfeld, J. H.: Secondary aerosol formation
25 from atmospheric reactions of aliphatic amines, *Atmos. Chem. Phys.*, 7, 2313–2337,
26 doi:10.5194/acp-7-2313-2007, 2007

27

28 Neubauer, K. R., Johnston, M. V., and Wexler, A. S.: Humidity effects on the mass
29 spectra of single aerosol particles, *Atmos. Environ.*, 32, 2521–2529, 1998.

30

31 O'Brien R.E., A. Laskin, J. Laskin, S. Liu, R. Weber, L. Russell, and A.H. Goldstein.
32 Molecular Characterization of Organic Aerosol Using Nanospray
33 Desorption/Electrospray Ionization Mass Spectrometry: CalNex 2010 field study.
34 *Atmospheric Environment* 68:265-272. doi:10.1016/j.atmosenv.2012.11.056, 2013

35

36 Pastor, S.H., J.O. Allen, L.S. Hughes, P. Bhave, G.R. Cass, and K.A. Prather,
37 Ambient single particle analysis in Riverside, California by aerosol time-of-flight mass
38 spectrometry during the SCOS97-NARSTO, *Atmospheric Environment*, 37, S239-
39 S258, 2003.

40

41 Pattanaik, S.; Huggins, F. E.; Huffman, G. P.; Linak, W. P.; Miller, C. A. XAFS studies
42 of nickel and sulfur speciation in residual oil fly-ash particulate matters (ROFA PM).
43 *Environ. Sci. Technol.* 41 (4), 1104–1110, 2007.



1

2 Pratt, K. A., Hatch, L. E., and Prather, K. A.: Seasonal volatility dependence of
3 ambient particle phase amines, *Environ. Sci. Technol.*, 43, 5276–5281, 2009.

4

5 Pratt, K. A. and Prather, K. A.: Mass spectrometry of atmospheric aerosols – Recent
6 developments and applications, Part II: On-line mass spectrometry techniques, *Mass
7 Spectrom. Rev.*, 31, 17–48, 2012

8

9 Prospero, J., Blades, E., Mathison, G., and Naidu, R.: Interhemispheric transport of
10 viable fungi and bacteria from Africa to the Caribbean with soil dust, *Aerobiologia*, 21,
11 1–19, 2005.

12

13 Qin, X. Y., Bhawe, P. V., and Prather, K. A.: Comparison of two methods for obtaining
14 quantitative mass concentrations from aerosol time-of-flight mass spectrometry
15 measurements, *Anal. Chem.*, 78, 6169–6178, 2006.

16

17 Rebotier, T. P. and Prather, K. A.: Aerosol time-of-flight mass spectrometry data
18 analysis: A benchmark of clustering algorithms, *Anal. Chim. Acta*, 585, 38–54, 2007.

19

20 Reche, C., Querol, X., Alastuey, A., Viana, M., Pey, J., Moreno, T., Rodríguez, S.,
21 González, Y., Fernández-Camacho, R., de la Rosa, J., Dall'Osto, M., Prévôt, A. S.
22 H., Hueglin, C., Harrison, R. M., and Quincey, P.: New considerations for PM, Black
23 Carbon and particle number concentration for air quality monitoring across different
24 European cities, *Atmos. Chem. Phys.*, 11, 6207–6227, doi:10.5194/acp-11-6207-
25 2011, 2011

26

27 Reche, C., Viana, M., Pandolfi, M., Alastuey, A., Moreno, T., Amato, F., Ripoll, A.,
28 and Querol, X.: Urban NH₃ levels and sources in a Mediterranean environment,
29 *Atmos. Environ.*, 57, 153–164, 2012.

30

31 Rehbein, P. J. G., Jeong, C.-H. , McGuire M. L., Yao X., Corbin J. C., Evans G. J.,
32 Cloud and fog processing enhanced gas-to-particle partitioning of trimethylamine.
33 *Environ. Sci. Technol.*, 45, 4346. doi:10.1021/ES1042113, 2011

34

35 Reilly, P. T. A.; Lazar, A. C.; Gieray, R. A.; Whitten, W. B.; Ramsey, J. M. The
36 elucidation of charge-transfer-induced matrix effects in environmental aerosols via
37 real-time aerosol mass spectral analysis of individual airborne particles. *Aerosol Sci.
38 Technol.*, 33, 135–152, 2000.

39



- 1 Reinard, M.S., Johnston, M.V., Ion formation mechanism in laser desorption
2 ionization of individual particles. *Journal of the American Society for Mass*
3 *Spectrometry* 19, 389e399, 2008
4
- 5 Salcedo, D., Onasch, T. B., Aiken, A. C., Williams, L. R., de Foy, B., Cubison, M. J.,
6 Worsnop, D. R., Molina, L. T., and Jimenez, J. L.: Determination of particulate lead
7 using aerosol mass spectrometry: MILAGRO/MCMA-2006 observations, *Atmos.*
8 *Chem. Phys.*, 10, 5371-5389, doi:10.5194/acp-10-5371-2010, 2010
9
- 10 Schoolcraft, T. A.; Constable, G. S.; Jackson, B.; Zhigilei, L. V.; Garrison, B. J.
11 Molecular dynamics simulations of laser disintegration of amorphous aerosol
12 particles with spatially non- uniform absorption. *Nucl. Instrum. Methods Phys. Res.*,
13 *Sect. B*, 180, 245-250, 2001.
14
- 15 Schofield, M.J., 2004. Sources and properties of airborne particulate matter. Ph.D.
16 Thesis. School of Geography, Earth and Environmental Sciences. Birmingham,
17 University of Birmingham.
18
- 19 Silva, P. J., Liu, D.-Y., Noble, C. A., and Prather, K. A.: Size and chemical
20 characterization of individual particles resulting from biomass burning of local
21 Southern California species, *Environ. Sci. Technol.*, 33, 3068–3076, 1999.
22
- 23 Silva, P. J. and Prather K. A.: Interpretation of mass spectra from organic
24 compounds in aerosol time-of-flight mass spectrometry, *Anal. Chem.*, 72, 3553-3562,
25 2000.
26
- 27 Silva, P. J.; Erupe, M. E.; Price, D.; Elias, J.; Malloy, Q. G. J.; Li, Q.; Warren, B.;
28 Cocker, D. R. Trimethylamine as precursor to secondary organic aerosol formation
29 via nitrate radical reaction in the atmosphere. *Environ. Sci. Technol*, 42 (13), 4689–
30 4696, 2008.
31
32
- 33 Shiraiwa, M., Kondo, Y., Moteki, N., Takegawa, N., Miyazaki, Y., and Blake, D. R.:
34 Evolution of mixing state of black carbon in polluted air from Tokyo, *Geophys. Res.*
35 *Lett.*, 34, L16803, doi:10.1029/2007GL029819, 2007.
36
- 37 Sleiman, M., Destailats, H., Smith, J., Liu, C.L., Ahmed, M., Wilson, K.R., Gundel,
38 L.A. Secondary organic aerosol formation from ozone-initiated reactions with nicotine
39 and secondhand smoke. *Atmospheric Environment* 44, 4191-4198, 2010.
40
- 41 Solazzo, E., Cai, X., and Vardoulakis, S.: Modelling wind flow and vehicle-induced
42 turbulence in urban streets, *Atmos. Environ.*, 42, 4918–4931, 2008.
43



- 1 Song, X. H., Hopke, P. K., Fergenson, D. P., and Prather, K. A.: Classification of
2 single particles analyzed by ATOFMS using an artificial neural network, ART-2A,
3 Anal. Chem., 71, 860–865, 1999
4
5
6 Spencer, M. T., Shields, L. G., Sodeman, D. A., Toner, S. M. and Prather K. A.:
7 Comparison of oil and fuel particle chemical signatures with particle emissions from
8 heavy and light duty vehicles, Atmos. Environ., 40, 5224-5235, 2006.
9
10 Su, Y. X., Sipin, M. F., Furutani, H., and Prather, K. A. Development and
11 Characterization of an Aerosol Time-of-Flight Mass Spectrometer with Increased
12 Detection Efficiency. Anal. Chem., 76(3):712–719, 2004.
13
14 Sureda, X., Martinez-Sanchez, J.M., Lopez, M.J., Fu, M., Agüero, F., Salto, E.,
15 Nebot, M.; Fernandez, E; Secondhand smoke levels in public building main
16 entrances: outdoor and indoor PM_{2.5} assessment, Tobacco Control 21:543–548,
17 2012.
18
19 Volkamer, R., Jimenez, J. L., San Martini, F., Dzepina, K., Zhang, Q., Salcedo, D.,
20 Molina, L. T., Worsnop, D. R. and Molina, M. J. Secondary organic aerosol formation
21 from anthropogenic air pollution: Rapid and higher than expected, Geophysical
22 Research Letters, 33, L17811, doi:10.1029/2006GL026899, 2006.
23
24 Wayne, R. P., Barnes, I., Biggs, P., Burrows, J. P., Canasa-Mas, C. E., Hjorth, J., Le
25 Bras, G., Moorgat, G. K., Perner, D., Poulet, G., Restelli, G., and Sidebottom, H.: The
26 nitrate radical physics, chemistry, and the atmosphere., Atmos. Environ., 25A, 1–203,
27 1991.
28
29 WHO: Health Aspects of Air pollution, Syst.Rew.of health asp. of air poll. in Europe,
30 Copenhagen, 2004
31
32 WHO (http://www.who.int/whosis/whostat/EN_WHS10_Full.pdf). Last access
33 December 2012
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Monitoring RS site	Particle type	particle number	%	Monitoring UB site	Particle type	particle number	%
Local & Regional EC	<i>EC Aged L</i>	277,151	31.1	Local & Regional EC	<i>EC Aged L</i>	52,074	23.8
	<i>EC Aged R</i>	202,227	22.7		<i>EC Aged R</i>	76,678	35.1
Secondary inorganic	<i>LRT-NIT</i>	102,909	11.6	Secondary inorganic	<i>LRT-NIT</i>	9,982	4.6
	<i>Loc-NIT</i>	37,790	4.2		<i>Loc-NIT</i>	15,931	7.3
	<i>LRT-SUL</i>	52,420	5.9		<i>LRT-SUL</i>	6,635	3.0
Sea Salt	<i>NaCl-NIT</i>	75,204	8.4	Sea Salt	<i>NaCl-NIT</i>	8,166	3.7
	<i>NaCl</i>	8,926	1.0		<i>NaCl</i>	31,853	14.6
Industry	<i>Fe</i>	7,201	0.8	Industry	<i>Fe</i>	5,564	2.5
	<i>Pb</i>	577	0.1		<i>Pb</i>	382	0.2
Combustion	<i>K-CN</i>	21,515	2.4	Combustion	<i>K-CN</i>	2,778	1.3
Monitoring RS site	Particle type	particle number	%	Monitoring UB site	Particle type	particle number	%
RS particles	<i>Amine (POA 58)</i>	6,698	0.8	UB particles	<i>Soil-Saharan</i>	2,842	1.3
	<i>Amine (SOA 114)</i>	3,672	0.4		<i>Soil-Ca</i>	2,482	0.9
	<i>Amine (SOA 59)</i>	2,141	0.2		<i>Oil-V</i>	1,875	0.9
	<i>Amine (EST 84)</i>	4,888	0.5		<i>Veg-KP</i>	3,897	1.7
	<i>Org. (Lub Oil)</i>	16,273	1.8				
	<i>Org. (OC-CHO)</i>	42,680	4.8				
	<i>Org. (Aro-NIT)</i>	15,306	1.7				
	<i>Org. (OC-NIT)</i>	13,295	1.5				
TOTAL RS		890,873	100	TOTAL UB		221,139	100

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6 Table 1 ATOFMS particle clusters identified from the SAPUSS campaign.

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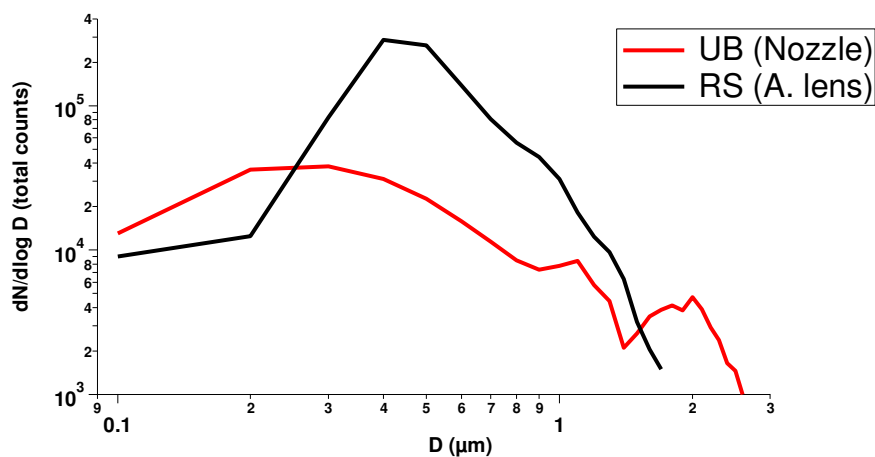
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5 Figure 1. Size distributions of collected ATOFMS particles at the two SAPUSS
6 monitoring sites

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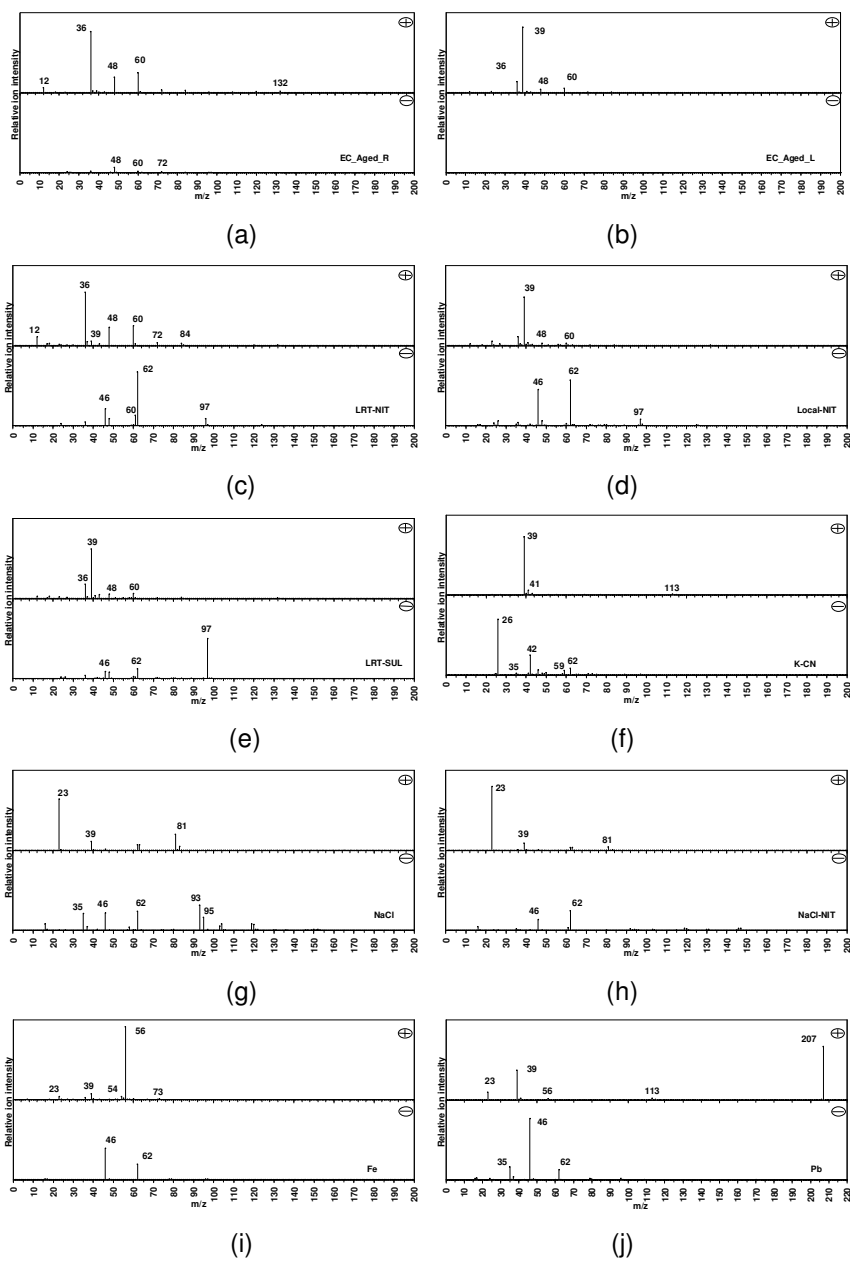
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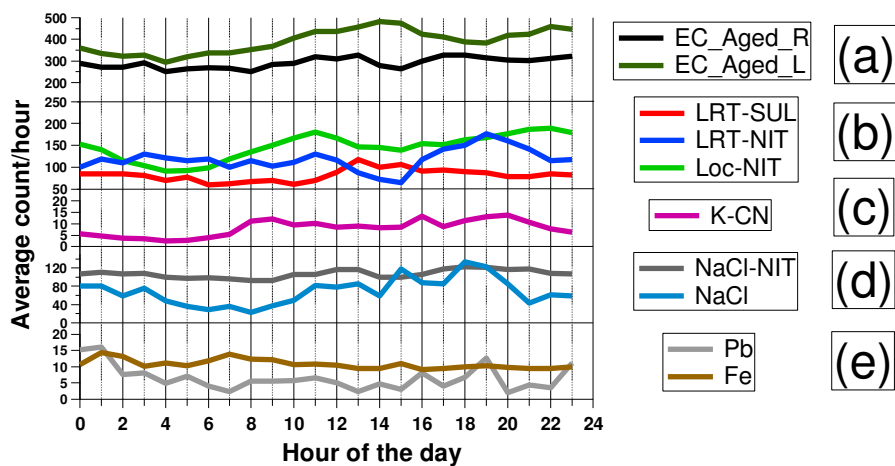


13 Figure 2. Average mass spectra of the 10 single particle types observed at both the
14 RS and UB sites.



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4 Figure 3. Diurnal profiles of the common particle types detected at both the RS and
5 UB sites. Differences between RS and UB sites were minor and average diurnal
6 profiles are presented.

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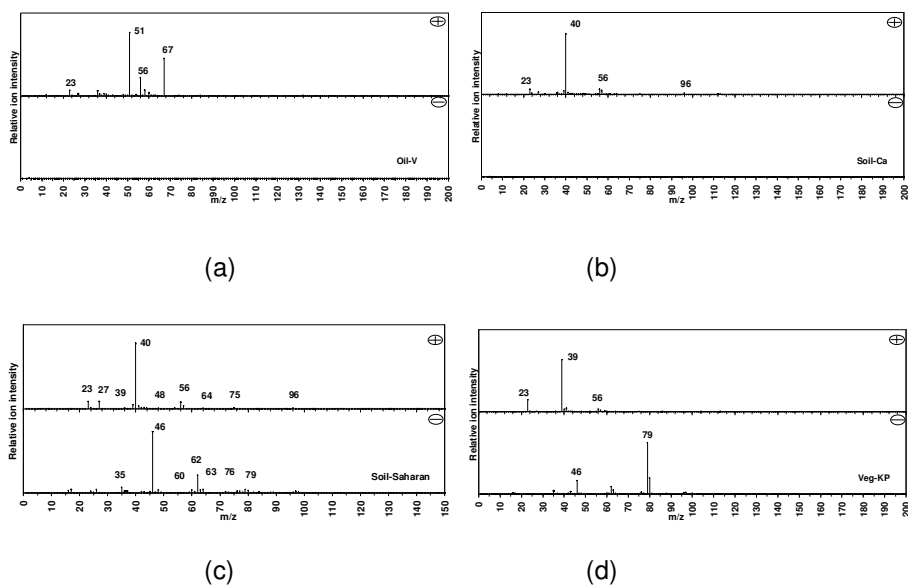
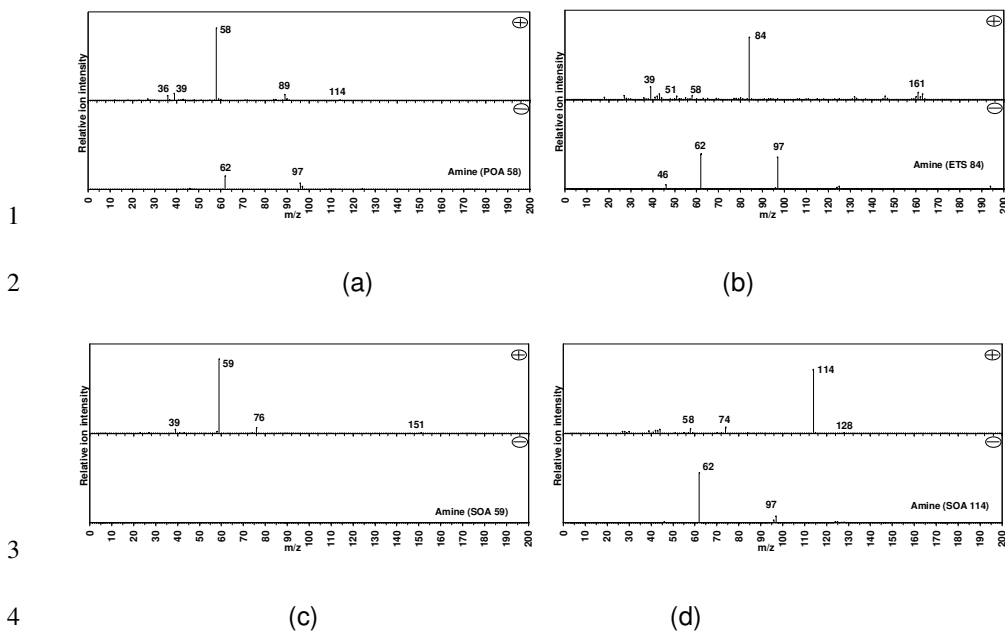


Figure 4. Average mass spectra of the 4 single particle types only observed at the

UB site.



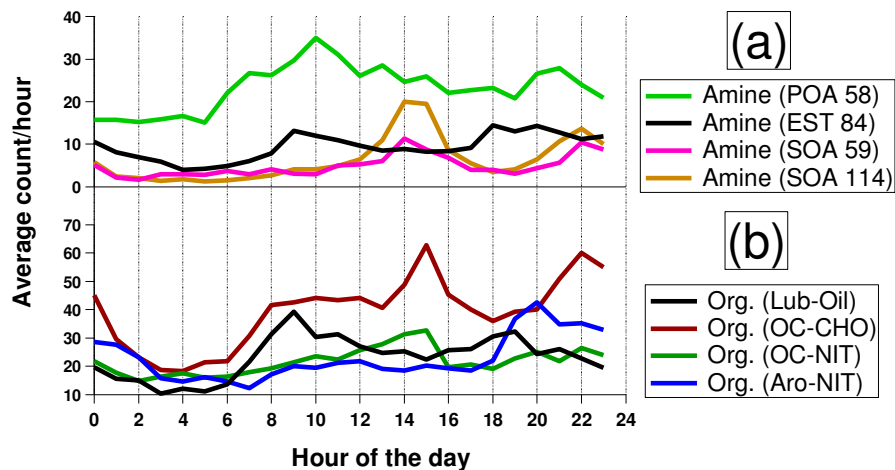
5 Figure 5. Average mass spectra of the 4 amine-containing single particle types only
6 observed at the RS site.

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4 Figure 6. Diurnal trend of Amines (a) and organic-rich (b) particle types detected at
5 the RS site. . Differences between RS and UB sites were minor and average diurnal
6 profiles are presented.

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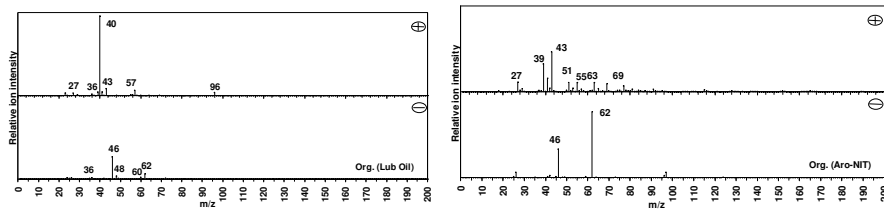
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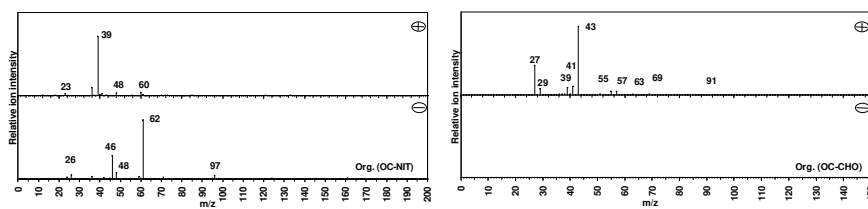


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(a)

(b)



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(c)

(d)

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8 Figure 7. Average mass spectra of the 4 OC-rich single particle types only observed
9 at the RS site.

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