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# Global carbon tetrachloride distributions obtained from the Atmospheric Chemistry Experiment (ACE)

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## Abstract

The first study of the global atmospheric distribution of carbon tetrachloride ( $\text{CCl}_4$ ), as a function of altitude and latitude, was performed using solar occultation measurements obtained by the Atmospheric Chemistry Experiment (ACE) mission using Fourier transform spectroscopy. A total of 8703 profile measurements were used in the study taken between February 2004 and August 2007. The zonal distribution of carbon tetrachloride displays a slight hemispheric asymmetry and decreasing concentration with increasing altitude at all latitudes. Maximum carbon tetrachloride concentrations are situated below 10 km in altitude with VMR (Volume Mixing Ratio) values of 100–130 ppt (parts per trillion). The highest concentrations are located about the equator and at mid-latitudes, particularly for latitudes in heavily industrialised regions (20–45° N), with values declining towards the poles. Global distributions obtained from ACE were compared with predictions from three chemistry transport models. The ACE dataset gives unique global and temporal coverage of carbon tetrachloride and its transport through the atmosphere. An estimated lifetime for carbon tetrachloride of  $34 \pm 5$  years was determined through correlation with CFC-11.

## 1 Introduction

The synthesis of carbon tetrachloride ( $\text{CCl}_4$ , also known as tetrachloromethane) from chloroform and chlorine was first reported by Regnault in 1839. Chemically inert, carbon tetrachloride was originally used as a fumigant, fire extinguisher and predominantly as a cleaning reagent or industrial solvent (Galbally, 1976; Altshuller, 1976). The chronic toxicity associated with carbon tetrachloride led to its replacement with less harmful solvents (Singh, 1976). Nevertheless with the growing consumption of chlorofluorocarbons (CFCs) in the 1950s carbon tetrachloride became an important feedstock in the production of CFC-11 and CFC-12 (Simmonds et al., 1998). Despite the reduction in the use of carbon tetrachloride it is still a major component of the

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atmospheric chlorine budget.

The main, if not sole, source of carbon tetrachloride is from anthropogenic emissions, although it has been suggested that there is a small biogenic component (Lovelock and Maggs, 1973; Butler et al., 1999). However attempts to reconstruct an atmospheric history have proven inconclusive (Sturrock et al., 2002). Due to its longevity in the atmosphere it is believed that the main sink mechanism for carbon tetrachloride is through photochemical destruction in the stratosphere (Galbally, 1976). The ocean (Lovelock and Maggs, 1973; Yvon-Lewis et al., 2002) and soil (Happell and Roche, 2003) are thought to be additional less significant sinks. The loss of carbon tetrachloride caused by ocean sinks reduced the estimated atmospheric lifetime of carbon tetrachloride to 26 years (World Meteorological Organisation, 2002).

Following concerns over the catalytic destruction of ozone by chlorine-containing species (Molina and Rowland, 1974) estimates of global emissions of carbon tetrachloride have been attempted (Galbally, 1976; Simmonds et al., 1988, 1998; Singh, 1976), although poor industrial records and an underestimation of emissions in Asia (Palmer et al., 2003) and other parts of the world have led to significant levels of uncertainty. The discovery of the ozone hole by the British Antarctic Survey (Farman et al., 1985) led to compounds with high ozone depleting potentials (ODPs), such as carbon tetrachloride, being phased out through the implementation of the Montreal Protocol (The Montreal Protocol on Substances that Deplete the Ozone Layer and its Amendments and Adjustments, 1987) and its amendments.

Carbon tetrachloride is a potent greenhouse gas with a global warming potential (GWP) of 1400, relative to CO<sub>2</sub> which has a GWP of 1, on a 100 year time horizon (Solomon et al., 2007). As carbon tetrachloride is a relatively short-lived chlorine-containing compound compared to CFCs, the reduction of approximately 1% a year since the early 1990s in surface level concentrations has had a considerable effect in reducing anthropogenic radiative forcing. Surprisingly the Montreal Protocol has already had a greater effect on climate protection than that of the first commitment period of the Kyoto Protocol (Velders et al., 2007).

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The first atmospheric measurements of carbon tetrachloride were taken by Lovelock and Maggs (1973) over the Atlantic. Since 1978 the Advanced Global Atmospheric Gases Experiment (AGAGE) and its predecessors have made continuous ground based in situ measurements (Prinn et al., 2000) from clean sea air at five locations. NASA's intensive airborne measurement campaigns over the Pacific, including Pacific Exploratory Mission (PEM) West A and B, as well as the Transport and Chemical Evolution over the Pacific (TRACE-P) in 1991, 1994 and 2001, respectively, sampled carbon tetrachloride evolution and transport at altitudes below 12 km (Blake et al., 1996, 2003). In situ gas chromatography (GC) measurements taken on board the ER-2 aircraft (Elkins et al., 1996; Romashkin et al., 2000) have compared carbon tetrachloride concentrations with reported GC measurements taken on seven balloon flights between 7° S and 67° N (Moore et al., 2003) from 1996 to 2000 and one MkIV balloon flight using FTS (Fourier transform spectroscopy) at 65° N (Toon et al., 1999). ATMOS (Atmospheric Trace MOlecule Spectroscopy) volume mixing ratio (VMR) profiles for carbon tetrachloride between 30° N and 51° N taken using a Fourier transform spectrometer (FTS) on the ATLAS-3 Space Shuttle Mission in November 1994 were also compared with ER-2 aircraft data (Chang et al., 1996). Stratospheric measurements were reported using Cryogenic Infrared Radiance Instrumentation (CIRRIS-1A) during a short space mission (Zhou et al., 1998). However there has been a lack of stratospheric and global measurements. To the best of our knowledge, there have been no prior reports of a global distribution of carbon tetrachloride using satellite data.

## 2 Observations and retrievals

The Atmospheric Chemistry Experiment (ACE), also referred to as SCISAT, is a Canadian satellite used for remote sensing of trace gases in the Earth's atmosphere from a low circular orbit (altitude 650 km, inclination 74°). The primary instrument on the ACE satellite is a high-resolution Fourier transform spectrometer (FTS) with coverage of the 750 to 4400 cm<sup>-1</sup> spectral region. The spectrometer has a 0.02 cm<sup>-1</sup> spectral reso-

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lution with a maximum optical path difference of  $\pm 25$  cm (Bernath et al., 2005). The ACE-FTS records spectra by solar occultation. The sun tracker locks onto the centre of the sun whilst a set of transmission spectra are taken through the limb of the Earth's atmosphere during sunrise and sunset. VMR profiles of trace species from the mid-troposphere through to the lower thermosphere, along with temperature and pressure profiles, are retrieved from the set of transmission spectra (Boone et al., 2005). Over 30 trace species are retrieved from ACE-FTS spectra, including multiple isotopologues for some species, with a vertical resolution of 3–4 km.

The current retrieval version for the ACE-FTS is version 2.2, however carbon tetrachloride was not reported in the official version 2.2 results. The data employed here are from a “research processing” dataset generated using the ACE-FTS analysis software. The retrieval window for carbon tetrachloride is centered at  $796.5\text{ cm}^{-1}$  and is  $18\text{ cm}^{-1}$  wide. The altitude range of the retrieval varies as a function of latitude, extending from 7 to 30 km at the equator and 6 to 25 km at the poles. Treating different isotopologues of a given molecule as separate interferers (with each isotopologue having a different VMR profile) there are a total of 13 interferers in the retrieval of carbon tetrachloride. These interferers are  $\text{H}_2\text{O}$  isotopologues  $\text{H}_2^{16}\text{O}$ ,  $\text{H}_2^{18}\text{O}$  and  $\text{H}_2^{17}\text{O}$ ,  $\text{CO}_2$  isotopologues  $^{12}\text{C}^{16}\text{O}_2$ ,  $^{13}\text{C}^{16}\text{O}_2$  and  $^{12}\text{C}^{16}\text{O}^{18}\text{O}$ ,  $\text{O}_3$ ,  $\text{HO}_2\text{NO}_2$ ,  $\text{ClONO}_2$ ,  $\text{C}_2\text{H}_6$ ,  $\text{C}_2\text{H}_2$ , and HCFC-22. Additional microwindows are used to improve the retrieval results for the following interferers:  $\text{H}_2\text{O}$  isotopologues  $\text{H}_2^{18}\text{O}$  and  $\text{H}_2^{17}\text{O}$ ,  $\text{CO}_2$  isotopologues  $^{13}\text{C}^{16}\text{O}_2$  and  $^{12}\text{C}^{16}\text{O}^{18}\text{O}$ ,  $\text{ClONO}_2$ ,  $\text{C}_2\text{H}_6$ ,  $\text{C}_2\text{H}_2$ , and HCFC-22. Spectroscopic parameters for the interferer  $\text{HO}_2\text{NO}_2$  were obtained using pseudo-lines from the ATMOS line list (Brown et al., 1996). Cross-section data for  $\text{ClONO}_2$ , HCFC-22, and  $\text{CCl}_4$ , along with line parameters for all the other interferers in the microwindow, were obtained from the HITRAN 2004 dataset (Rothman et al., 2005). There is clear evidence of line mixing in a  $\text{CO}_2$  Q-branch contained in the  $\text{CCl}_4$  microwindow (Kochel et al., 1997), but this is not included in the forward model calculation and is simply neglected. For a tropical occultation ss11613 at 8.3 km a plot of the species which contribute to the window in which carbon tetrachloride is retrieved are shown (Fig. 1). In Fig. 1  $\text{ClONO}_2$  is weak

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and has a much stronger absorption in the stratosphere and in the polar regions.

In polar winter and spring the VMRs of chemical species such as carbon tetrachloride are low as a result of descent within the polar vortex. Occultation measurements taken by ACE inside the polar vortex or at the vortex edge have therefore been excluded. The 10978 occultation measurements taken by ACE between February 2004 and August 2007 were filtered using the derived meteorological products classified using potential vorticity values (Manney et al., 2007) obtained from Met Office data (Swinbank and O'Neill, 1994; Swinbank et al., 2002; Davies et al., 2005), in a similar approach to that adopted by Nassar et al. (2005) 2275 profiles were classified as within or on the edge of the polar vortex and were discarded, while 8703 were classified as extravortex occultations. The geographic distribution of ACE sampling is shown in Fig. 2.

### 3 Results and discussion

Data starting at 90° S and ending at 90° N were separated into 36 equal latitude zones with a width of five degrees; six bins (90° S–70° S and 80° N–90° N) were removed in the polar regions due to the small number of extravortex occultations. An average profile was calculated for each latitude bin in order to improve the precision of the global distribution. Carbon tetrachloride has a lifetime estimated to be approximately 26 years (WMO report, 2006) and ground based measurements show a slight seasonal cycle (Prinn et al., 2000). The 2006 WMO report estimates that the surface concentration of carbon tetrachloride is declining at a rate of  $1 \pm 0.1$  ppt year<sup>-1</sup>. ACE measurements have been averaged over the February 2004 to August 2007 time period to obtain a global distribution. A contour plot of the global distribution was created using the VMR data for the 30 averaged bins spanning the latitudes 70° S to 80° N between the altitudes of 8 and 30 km. The data for the 30 average bins is provided in supplementary Table S1 <http://www.atmos-chem-phys-discuss.net/9/13299/2009/acpd-9-13299-2009-supplement.pdf>.

The global distribution for carbon tetrachloride, shown in Fig. 3 as a contour map,

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is nearly symmetric about the equator. There is a general trend that the VMR monotonically decreases as a function of altitude; concentrations decline towards the poles. Carbon tetrachloride is emitted into the troposphere, and this is where the highest VMR values between 100 and 130 ppt are observed. The greatest concentrations corresponded to latitudes that encompass major industrial regions (20–50° N). Lower stratospheric concentrations are significantly higher between 20° S and 20° N compared to higher latitudes where mixing ratios are 20–40 ppt lower. This pattern can be attributed to the Brewer-Dobson circulation and the general appearance is typical of long lived gases such as CFCs that have surface sources and are destroyed in the stratosphere. In the upper stratosphere there is a decline in the VMR of carbon tetrachloride, which becomes negligible above 25 km in the polar latitudes and 28 km in the latitudes near the equator, corresponding to altitudes close to the detection limit. The drop in concentration is consistent with transport and photochemical destruction in the stratosphere. There is some asymmetry at higher latitudes with lower mixing ratios observed at southern polar latitudes compared to the equivalent northern latitudes.

To identify hemispheric differences in VMR, the extravortex occultations were divided into seven different latitude zones (60–90° S, 45–60° S, 20–45° S, 20° S–20° N, 20–45° N, 45–60° N and 60–90° N) and an average VMR for each zone was determined (supplementary Table S2 <http://www.atmos-chem-phys-discuss.net/9/13299/2009/acpd-9-13299-2009-supplement.pdf>). The VMR average for each zone was then plotted as a function of altitude, as shown in Fig. 4. In the mid and lower stratosphere carbon tetrachloride VMRs show little difference between the corresponding southern and northern hemispheric zones. Any differences may be attributed to the severity of the winter polar vortex in the Southern Hemisphere. The greatest decline in the concentration of carbon tetrachloride, as a function of altitude, is observed in the zones closest to the poles (60–90° S and 60–90° N). Above the troposphere concentrations are greatest in the equatorial zone (20° S–20° N) at all altitudes. Tropospheric VMR values are higher between 20° N and 60° N consistent with emission predominantly in the Northern Hemisphere. The overall pattern is consistent with carbon tetrachloride

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having a long chemical lifetime with destruction in the upper stratosphere.

The majority of carbon tetrachloride measurements have been taken at ground level or from aircraft in the troposphere and the lower stratosphere. Prinn et al. (2000) reported carbon tetrachloride concentrations at ground level had declined to approximately 97–100 ppt by June 1998 at the five global AGAGE sites. However a reconstructed history of carbon tetrachloride up to the year 2000 from polar firn (unconsolidated snow) air estimated concentrations to be over 100 ppt (Butler et al., 1999; Martinerie et al., 2009). Comparison of PEM West B and TRACE-P aircraft campaign data showed a decrease in the VMR of carbon tetrachloride over the Pacific from 107–108 ppt in 1994 to 100–99 ppt in 2001. The 2006 WMO Report cited the global average concentration of carbon tetrachloride as 94.6 ppt in 2004, dropping at a rate of approximately  $1.0 \text{ ppt yr}^{-1}$ . Our remote sensing observations show considerable variation in tropospheric concentrations at 8 km from about 100 ppt near the poles to as high as 130 ppt at northern mid-latitudes. These high tropospheric values and ACE data errors are discussed below.

Balloon measurements allow comparison of vertical profiles with ACE data over a wide range of altitudes, including measurements in the upper stratosphere. The JPL MkIV interferometer (Toon, 1991) is a balloon-borne solar occultation FTS, similar to ACE. The MkIV has performed 21 balloon flights since 1989, each covering altitudes from the cloud tops to the balloon (35–40 km) at 2–4 km vertical resolution. The MkIV  $\text{CCl}_4$  retrievals use the same absorption band as ACE, and also ignore line mixing from the  $\text{CO}_2$  Q-branch. In Figs. 5, 6 and 7 MkIV balloon FTS measurements are displayed (Toon et al., 1999). These were taken during a number of campaigns at Ft. Sumner, New Mexico (33–35° N), Fairbanks, Alaska (64–70° N) and Esrange, Sweden (62–70° N). Carbon tetrachloride profiles at the three locations were compared with the average ACE VMR profile of the latitude bin corresponding to the latitude of the balloon flights (Figs. 5–7).

Measurements obtained from balloon launches in Ft. Sumner, New Mexico (Fig. 5) have very similar VMR profiles compared with ACE measurements and generally agree

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within measurement error. The New Mexico profiles decline at a similar rate to the average ACE profiles from the appropriate latitude bin. The majority of MkIV measurements cease at 11 km, with VMRs of 120 ppt in agreement with values from ACE (Fig. 3). In Fig. 6 all of the MkIV balloon profiles show good agreement with ACE retrievals in the stratosphere but there is a large discrepancy at tropospheric altitudes for the ascent and descent profiles measured on 8 July 1997.

The magnitude of systematic errors at low altitude is greatest using solar occultation technique. The measurements from Erange (Fig. 7) have VMRs that decrease more rapidly at stratospheric altitudes compared with ACE and other MkIV profiles. The more rapid decline in the lower stratospheric VMR of carbon tetrachloride in the balloon profiles can be attributed to the fact that the measurements were carried out in spring and are influenced by descent in the polar vortex, whilst the ACE zonal means are an average of extravortex profiles. In general there is good agreement between the MkIV balloon profiles and ACE data, especially in the stratosphere however in the lower troposphere there is considerable variation.

The global distribution of carbon tetrachloride has been compared to calculations from three atmospheric models: The AER 2-D model (Weisenstein et al., 2004), the GMI 3-D chemical transport model (CTM) (Strahan et al. 2007) and the SLIMCAT 3-D CTM (Chipperfield 1999; Feng et al., 2007). All of the model runs used surface carbon tetrachloride mixing ratios from WMO (2006) as the lower boundary condition. These mixing ratios decrease, for example, from 98.7 ppt in January 2000 to 93.9 ppt in January 2005. In the models carbon tetrachloride has a very long-lifetime in the troposphere and the imposed surface mixing ratios are well-mixed. In the stratosphere the models destroy carbon tetrachloride by photolysis and by reaction with  $O(^1D)$ . The predicted stratospheric distribution therefore depends to a large extent on the model stratospheric circulation. Stratospheric models are known to exhibit large differences with respect to each other and to observations for this circulation. In particular, in the past models have calculated a circulation which is too strong, leading to an age-of-air which is too young. For the runs shown here the GMI model (horizontal resolution

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2°×2.5° and 42 levels from surface to 0.0158 hPa) was forced with six-hourly averaged GEOS4DAS winds. These produce a much better circulation than previous instantaneous GEOS4 winds. The SLIMCAT CTM run (horizontal resolution 5.6°×5.6° and 32 levels from surface to ~0.1 hPa) was forced with European Centre for Medium Range Weather Forecasts analyses and used an isentropic vertical coordinate. Vertical transport in the stratosphere was diagnosed from calculated heating rates. This approach also gives a good representation of the stratospheric age-of-air (see Chipperfield 2006). The AER 2-D model (horizontal resolution 9.5° and vertical resolution of 1.2 km from the surface to 0.02 hPa) used a climatological residual circulation from Fleming et al. (1999) but temperature and aerosol surface area matching the model year.

Figure 8 shows the zonal mean distribution of carbon tetrachloride calculated by the three models averaged over the period of ACE observations. All three models show the same basic distribution with a well-mixed troposphere and carbon tetrachloride decreasing with height in the stratosphere. All three models show some hemispheric asymmetry, with higher concentrations in the Northern Hemisphere. The SLIMCAT model shows the greatest hemispheric difference which becomes more prominent towards the polar regions. The AER and GMI models are very comparable, as they have very similarly shaped distributions but the SLIMCAT model shows a much steeper decline in carbon tetrachloride outside of the tropics. Overall there is good general agreement for the latitudinal distribution between the models and the ACE data. However, at lower altitudes model concentrations are no higher than 90–100 ppt (a result of the model boundary conditions) and model concentrations are then lower throughout the stratosphere below 30 km.

Figure 9 compares the altitude profile of the averaged ACE carbon tetrachloride profile in two latitude bins with corresponding averaged model data. Focusing on areas where ACE data shows high concentrations, such as the tropics (5° S–5° N) and northern mid-latitudes (30–40° N), similarities can be identified. In the 5° S–5° N tropical region (Fig. 9a) all models show good agreement with ACE data in the lower strato-

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sphere in terms of the vertical gradient of carbon tetrachloride. However, the models differ considerably in the troposphere below 15 km. A similar effect is seen at mid latitudes 30–40° N (Fig. 9b) with the models having similar vertical gradients between 15 and 25 km but a tropospheric mixing ratio constrained to lower values than observed by ACE.

ACE retrievals appear to give an overestimate of the carbon tetrachloride concentration in the troposphere as compared to surface values (e.g. Montzka et al., 1999) due to the inherent difficulties in low altitude retrievals for limb viewing instruments. At low altitude the spectra become very complex and there are significant, systematic features in the fitting residuals that grow larger as you move to lower altitude. These include large residuals from H<sub>2</sub>O lines in the spectral window (Boone et al., 2007) and residuals from neglecting line-mixing for the CO<sub>2</sub> Q-branch in the window. However, the magnitude of these errors is difficult to estimate. The statistical errors in the retrieval below 15 km remain low at the typical value of 5%, but the systematic errors can grow to 20–30%. The systematic error in the cross section data (Nemtchinov and Varanasi, 2003) used in our retrievals is estimated to be as large as 10%.

The stratospheric lifetime of carbon tetrachloride can be estimated using tracer-tracer correlations with species of known lifetimes (Plumb and Ko, 1992). Using Eq. (1) and the method reported by Barkley et al. (2008) for OCS from ACE, the stratospheric lifetime of carbon tetrachloride was estimated by correlation with CFC-11 and CFC-12. Coincident CFC-11 and CFC-12 measurements were used because their stratospheric lifetimes are well known (Volk et al., 1997). The plot of carbon tetrachloride against the tracer should provide a linear correlation between the two gases in the lower stratosphere (Fig. 13). From this correlation if the lifetime is known for one of the gases, then the lifetime of the other species can be estimated (Plumb and Ko, 1992) using Eq. (1),

$$\frac{\tau_1}{\tau_2} = \frac{d\sigma_2}{d\sigma_1} \frac{\sigma_1}{\sigma_2}, \quad (1)$$

where  $\tau_1$  and  $\tau_2$  are the lifetimes and  $\sigma_1$  and  $\sigma_2$  the VMRs. The slope of the correlation is  $d\sigma_2/d\sigma_1$ . The steady state lifetimes used in Eq. (1) for CFC-11 and CFC-12 were

45±7 and 87±17 years, respectively (Volk et al., 1997). The estimated lifetime given for carbon tetrachloride by plotting against CFC-11 was approximately 34±5 years whilst plotting carbon tetrachloride against CFC-12 gave a value close to 32±7 years for the lifetime. The correlation with CFC-11 (Fig. 10) is much more satisfactory than for CFC-12 which shows more curvature; we therefore place more confidence in the value obtained using CFC-11. These estimates are similar to numbers reported in previously published work. Volk et al. (1997) estimated the steady state lifetime of carbon tetrachloride as 32±6 years. A similar correlation plot between CFC-11 and carbon tetrachloride outputs from the SLIMCAT model gave a lifetime of 38±5 years (not shown).

#### 4 Summary and conclusion

The first near global distribution of carbon tetrachloride has been obtained from 8703 ACE occultation measurements covering altitudes from 8 to 30 km over 30 latitude bins covering 70° S–80° N. Low altitude carbon tetrachloride VMRs are estimated to be between 100 and 130 ppt, somewhat higher than in situ observations. Tropospheric VMRs were highest at latitudes corresponding to industrialised regions (20–45° N), and near the equator.

There appears to be only a small hemispheric asymmetry, except for the polar regions where the VMR is significantly lower in the Antarctic. There is good general agreement between model predictions and experimental data. The vertical gradient of the model VMR profiles agree well with ACE data in the lower stratosphere (15–25 km), although standard halocarbon scenarios used to constrain stratospheric models (e.g. WMO 2007) give boundary conditions smaller than retrieved from ACE. The stratospheric lifetime of carbon tetrachloride is estimated to be 34±5 years from the carbon tetrachloride-CFC-11 tracer-tracer correlation.

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5 **References**

- Altshuller, A. P.: Average Tropospheric Concentration of Carbon Tetrachloride Based on Industrial Production, Usage, and Emissions, *Environ. Sci. Technol.*, 10, 596–598, 1976.
- Barkley, M. P., Palmer, P. I., Boone, C. P., Bernath, P. F., and Suntharalingam, P.: Global distributions of carbonyl sulfide in the upper troposphere and stratosphere, *Geophys. Res. Lett.*, 35, L14810, doi:10.1029/2008GL034270, 2008.
- 10 Bernath, P. F., McElroy, C. T., Abrams, M. C., et al.: Atmospheric Chemistry Experiment (ACE): Mission Overview, *Geophys. Res. Lett.*, 32, L15S01, doi:10.1029/2005GL022386, 2005.
- Blake, D. R., Chen, T.-Y., Smith Jr., T. W., Wang, C. J.-L., Wingenter, O. W., Blake, N. J., Rowland, F. S., and Meyer E. W.: Three-dimensional distribution of nonmethane hydrocarbons and halocarbons over the northwestern Pacific during the 1991 Pacific Exploratory Mission (PEM-West A), *J. Geophys. Res.*, 101, 1763–1778, 1996.
- 15 Blake, N. J., Blake, D. R., Simpson, I. J., et al.: NMHCs and halocarbons in Asian continental outflow during the Transport and Chemical Evolution over the Pacific (TRACE-P) Field Campaign: Comparison With PEM-West B, *J. Geophys. Res.*, 108, 8806, doi:10.1029/2002JD003367, 2003.
- Boone, C. D., Nassar, R., Walker, K. A., Rochon, Y., McLeod, S. D., Rinsland, C. P., and Bernath, P. F.: Retrievals for the Atmospheric Chemistry Experiment Fourier Transform Spectrometer, *Appl. Opt.*, 44, 7218–7231, 2005.
- Boone, C. D., Walker, K. A., and Bernath, P. F.: Speed-dependent Voigt profile for water vapour in infrared remote sensing applications, *J. Quant. Spectrosc. Radiat. Transfer*, 105, 525–532, 2007.
- 25 Brown, L. R., Gunson, M. R., Toth, R. A., Irion, F. W., Rinsland, C. P., and Goldman, A.: 1995 Atmospheric Trace Molecule Spectroscopy (ATMOS) linelist, *Appl. Opt.*, 35, 2828–2848, 1996.
- 30 Butler, J. H., Battle, M., Bender, M. L., Montzka, S. A., Clarke, A. D., Saltzman, E. S., Sucher,

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- C. M., Severinghaus, J. P., and Elkins, J. W.: A record of atmospheric halocarbons during the twentieth century from polar firn air, *Nature*, 399, 749–755, 1999.
- Chang, A. Y., Salawitch, R. J., Michelsen, H. A., et al.: A comparison of measurements from ATMOS and instruments aboard the ER-2 aircraft: Halogenated gases, *Geophys. Res. Lett.*, 23, 2393–2396, 1996.
- 5 Chipperfield, M. P.: Multiannual Simulations with a Three-Dimensional Chemical Transport Model, *J. Geophys. Res.*, 104, 1781–1805, 1999.
- Chipperfield, M. P.: New Version of the TOMCAT/SLIMCAT Off-Line Chemical Transport Model: Intercomparison of Stratospheric Tracer Experiments, *Q. J. R. Meteorol. Soc.*, 132, 1179–1203, doi:10.1256/qj.05.51, 2006.
- 10 Davies, T., Cullen, M. J. P., Malcolm, A. J., Mawson, M. H., Staniforth, A., White, A. A., and Wood, N.: A new dynamical core for the Met Office's global and regional modelling of the atmosphere, *Q. J. R. Meteorol. Soc.*, 131, 1759–1782, 2005.
- Farman, J. C., Gardiner, B. G., and Shanklin, J. D.: Large losses of total ozone in Antarctica reveal seasonal ClO<sub>x</sub>/NO<sub>x</sub> interaction, *Nature*, 315, 207–210, 1985.
- 15 Feng, W., Chipperfield, M. P., Dorf, M., Pfeilsticker, K., and Ricaud, P.: Mid-latitude Ozone Changes: Studies with a 3-D CTM Forced by ERA-40 Analyses, *Atmos. Chem. Phys.*, 7, 2357–2369, 2007, <http://www.atmos-chem-phys.net/7/2357/2007/>.
- 20 Fleming, E. L., Jackman, C. H., Stolarski, R. S., and Considine, D. B.: Simulation of stratospheric tracers using an improved empirically-based two-dimensional model transport formulation, *J. Geophys. Res.*, 104, 23911–23934, 1999.
- Galbally, I. E.: Man-Made Carbon Tetrachloride in the Atmosphere, *Science*, 193, 573–576, 1976.
- 25 Happell, J. D. and Roche, M. P.: Soils: A global sink of atmospheric carbon tetrachloride, *Geophys. Res. Lett.*, 30, 1088, doi:10.1029/2002GL015957, 2003.
- Kochel, J.-M., Hartmann, J.-M., Camy-Peyret, C., Rodrigues, R., and Payan, S.: Influence of line mixing on absorption by CO<sub>2</sub> Q branches in atmospheric balloon-borne spectra near 13 μm, *J. Geophys. Res.*, 102, 12891–12899, 1997.
- 30 Lovelock J. E. and Maggs R. J.: Halogenated Hydrocarbons in and over the Atlantic, *Nature*, 241, 194–196, 1973.
- Manney, G. L., Daffer, W. H., Zawodny, J. M., Bernath, P. F., Hoppel, K. W., Walker, K. A., Knosp, B.W., Boone, C., Remsberg, E. E., Santee, M. L., Harvey, V. L., Pawson, S., Jackson,

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D. R., Deaver, L., McElroy, C. T., McLinden, C. A., Drummond, J. R., Pumphrey, H. C., Lambert, A., Schwartz, M. J., Froidevaux, L., McLeod, S., Takacs, L. L., Suarez, M. J., Trepte, C. R., Cuddy, C. D., Livesey, N. J., Harwood, R. S., and Waters, J. W.: Solar occultation satellite data and derived meteorological products: Sampling issues and comparisons with Aura Microwave Limb Sounder, *J. Geophys. Res.*, 112, D24S50, doi:10.1029/2007JD008709, 2007.

Martinerie, P., Nourtier-Mazauric, E., Barnola, J.-M., Sturges, W. T., Worton, D. R., Atlas, E., Gohar, L. K., Shine, K. P., and Brasseur, G. P.: Long-lived halocarbon trends and budgets from atmospheric chemistry modelling constrained with measurements in polar firn, *Atmos. Chem. Phys., Discuss.*, 9, 1703–1726, 2009.

Molina, M. J. and Rowland, F. S.: Stratospheric sink for chlorofluoromethanes: chlorine atom-catalysed destruction of ozone, *Nature*, 249, 810–812, 1974.

The Montreal Protocol on Substances that Deplete the Ozone Layer and its Amendments and Adjustments, 1987.

Moore, F. L., Elkins, J. W., Ray, E. A., et al.: Balloonborne in situ gas chromatograph for measurements in the troposphere and stratosphere, *J. Geophys. Res.*, 108, 8330, doi:10.1029/2001JD000891, 2003.

Montzka, S. A., Butler, J. H., Elkins, J. W., Thompson, T. M., Clarke, A. D., and Lock, L. T.: Present and future trends in the atmospheric burden of ozone-depleting halogens, *Nature*, 398, 690–694, 1999.

Nassar, R., Bernath, P. F., Boone, C. D., Manney, G. L., McLeod, S. D., Rinsland, C. P., Skelton, R., and Walker, K. A.: ACE-FTS measurements across the edge of the winter 2004 Arctic vortex, *Geophys. Res. Lett.*, 32, L15S05, doi:10.1029/2005GL022671, 2005.

Nemtchinov, V. and Varanasi, P.: Thermal infrared absorption cross-sections of  $\text{CCl}_4$  needed for atmospheric remote sensing, *J. Quant. Spectrosc. Radiat. Transfer*, 82, 473–481, 2003.

Palmer, P. I., Jacob, D. J., Mickley, L. J., Blake, D. R., Sachse, G. W., Fuelberg, H. E., and Kiley, C. M.: Eastern Asian emissions of anthropogenic halocarbons deduced from aircraft concentration data, *J. Geophys. Res.*, 108, 4753, doi:10.1029/2003JD003591, 2003.

Prinn, R. G., et al.: A history of chemically and radiatively important gases in air deduced from ALE/GAGE/AGAGE, *J. Geophys. Res.*, 105, 17751–17792, 2000.

Plumb, R. A. and Ko, M. K. W.: Interrelationships between mixing ratios of long-lived stratospheric constituents, *J. Geophys. Res.*, 97, 10145–10156, 1992.

Romashkin, P. A., Hurst, D. F., Elkins, J. W., Dutton, G. S., Fahey, D. W., Dunn, R. E., Moore, F. L., Myers, R. C., and Hall, B. D.: In Situ Measurements of Long-Lived Trace Gases in the

Lower Stratosphere by Gas Chromatography, *J. Atmos. Oceanic Technol.*, 18, 1195–1204, 2000.

Rothman, L. S., Jacquemarta, D., Barbe, A., et al.: The HITRAN 2004 molecular spectroscopic database, *J. Quant. Spectrosc. Radiat. Transfer*, 96, 139–204, 2005.

5 Simmonds, P. G., Cunnold, D. M., Alyea, F. N., Cardelino, C. A., Crawford, A. J., Prinn, R. G., Fraser, P. J., Rasmussen, R. A., and Rosen, R. D.: Carbon Tetrachloride Lifetimes and Emissions Determined from Daily Global Measurements During 1978–1985, *J. Atmos. Chem.*, 7, 35–58, 1988.

10 Simmonds, P. G., Cunnold, D. M., Weiss, R. F., Miller, B. R., Prinn, R. G., Fraser, P. J., McCulloch, A., Alyea, F. N., and O'Doherty, S.: Global trends and emission estimates of CCl<sub>4</sub> from in situ background observations from July 1978 to June 1996, *J. Geophys. Res.*, 103, 16017–16027, 1998.

Singh, H. B., Fowler, D. P., and Peyton, T. O.: Atmospheric Carbon Tetrachloride: Another Man-Made Pollutant, *Science*, 192, 1231–1234, 1976.

15 Solomon, S., Qin, D., Manning, M., Alley, R. B., Bernsten, T., Bindoff, N. L., Chen, Z., Chidthaisong, A., Gregory, J. M., Hegerl, G. C., Heimann, M., Hewitson, B., Hoskins, B. J., Joos, F., Jouzel, J., Kattsov, V., Lohmann, U., Matsuno, T., Molina, M., Nicholls, N., Overpeck, J., Raga, G., Ramaswamy, V., Ren, J., Rusticucci, M., Somerville, R., Stocker, T.F., Whetton, P., Wood, R. A., and Wratt, D.: Technical Summary, In: *Climate Change 2007: The Physical Science Basis. Contribution of Working Group I to the Fourth Assessment Report of the Intergovernmental Panel on Climate Change*, edited by: Solomon, S., Qin, D., Manning, M., Chen, Z., Marquis, M., Averyt, K. B., Tignor, M., and Miller, H. L., Cambridge University Press, Cambridge, United Kingdom and New York, NY, USA, 2007.

20 Strahan, S. E., Duncan, B. N., and Hoor, P.: Observationally derived transport diagnostics for the lowermost stratosphere and their application to the GMI chemistry and transport model, *Atmos. Chem. Phys.*, 7, 2435–2445, 2007, <http://www.atmos-chem-phys.net/7/2435/2007/>.

25 Sturrock, G. A., Etheridge, D. M., Trudinger, C. M., Fraser, P. J., and Smith, A. M.: Atmospheric histories of halocarbons from analysis of Antarctic firn air: Major Montreal Protocol species, *J. Geophys. Res.*, 107, 4765, doi:10.1029/2002JD002548, 2002.

30 Swinbank, R. and O'Neill, A.: A stratosphere-troposphere data assimilation system, *Mon. Weather Rev.*, 122, 686–702, 1994.

Swinbank, R., Ingleby, N. B., Boorman, P. M., and Renshaw, R. J.: A 3-D variational data as-

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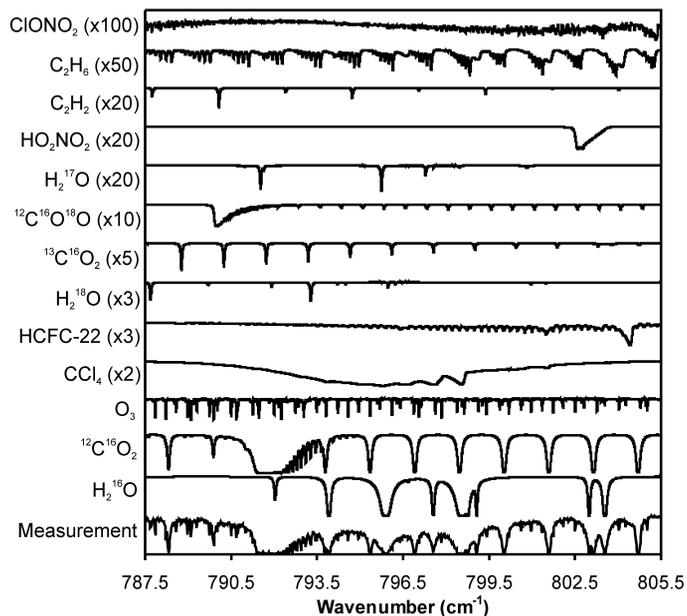
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- simulation system for the stratosphere and troposphere, Tech. Rep. 71, Met Office Numerical Weather Prediction Forecasting Research Scientific Paper, 2002.
- Toon, G. C.: The JPL MkIV Interferometer, Optics and Photonics News, 2, 19–21, 1991.
- Toon, G. C., Blavier, J. F., Sen, B., et al.: Comparison of MkIV balloon and ER-2 aircraft measurements of atmospheric trace gases, J. Geophys. Res., 104, 26779–26790, 1999.
- 5 Velders, G. J. M., Andersen, S. O., Daniel, J. S., Fahey, D. W., and McFarland, M.: The importance of the Montreal Protocol in protecting climate, PNAS, 104, 4814–4819, 2007.
- Volk, C. M., Elkins, J. W., Fahey, D. W., Dutton, G. S., Gilligan, J. M., Loewenstein, M., Podolske, J. R., Chance, K. R., and Gunson, M. R.: Evaluation of source gas lifetimes from stratospheric observations, J. Geophys. Res., 102, 25543–25564, 1997.
- 10 Weisenstein, D. K., Eluszkiewicz, J., Ko, M. K. W., Scott, C. J., Jackman, C. H., Fleming, E. L., Conside, D. B., Kinnison, D. E., Connell, P. S., and Rotman, D. A.: Separating chemistry and transport effects in 2-D models, J. Geophys. Res., 109, D18310, doi:10.1029/2004JD004744, 2004.
- 15 World Meteorological Organization (WMO), Scientific Assessment of Ozone Depletion: 2002, Geneva, 2003.
- World Meteorological Organisation (WMO): Scientific Assessment of Ozone Depletion: 2006, Geneva, Switzerland, 2007.
- Yvon-Lewis, S. A. and Butler, J. H.: The effect of oceanic uptake on the atmospheric lifetime of selected trace gases, J. Geophys. Res., 107, 4414, doi:10.1029/2001JD001267, 2002.
- 20 Zhou, D. K., Bingham, G. E., Anderson, G. P., and Nadile, R. M.: CIRRIIS-1A measurements of stratospheric carbon tetrachloride (CCl<sub>4</sub>) and carbon tetrafluoride (CF<sub>4</sub>), Geophys. Res. Lett., 25, 325–328, 1998.

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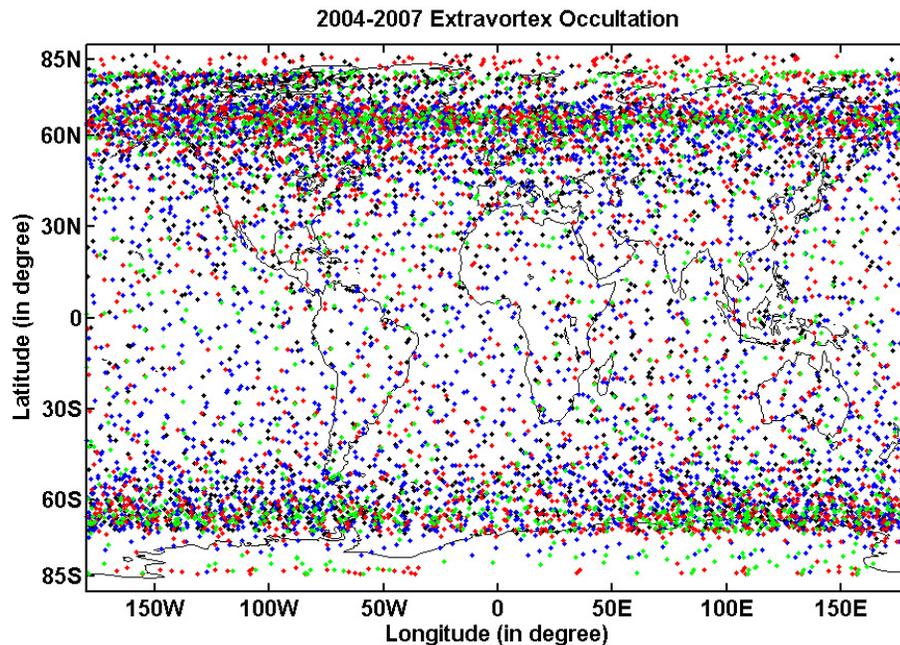


**Fig. 1.** The species which contribute to the spectral window in which carbon tetrachloride is retrieved. Weaker absorbers have been magnified and the baseline is offset in the vertical direction for clarity. The plot corresponds to the measurement (bottom panel) at a tangent height of 8.3 km in tropical occultation ss11613.

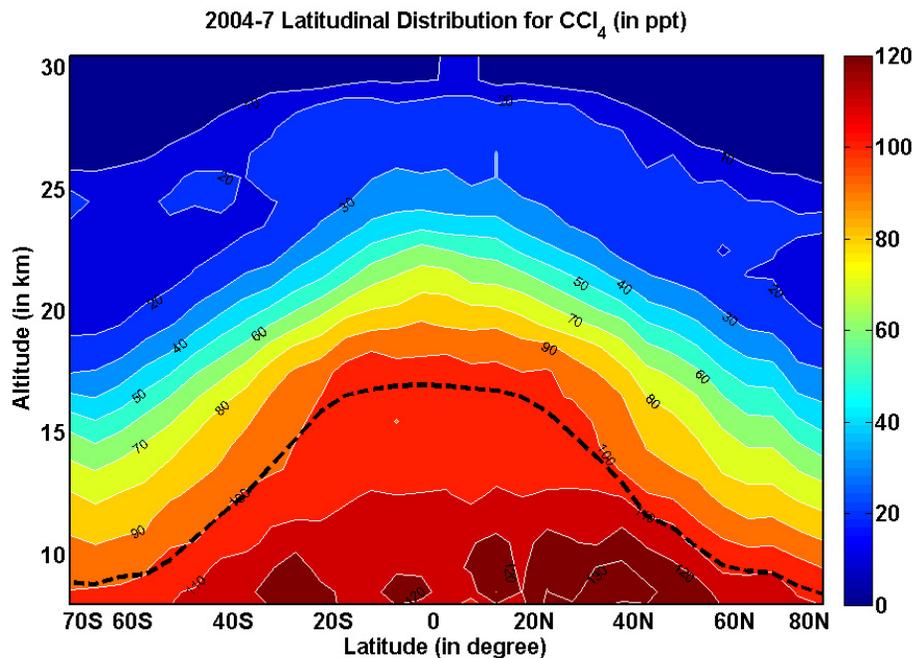
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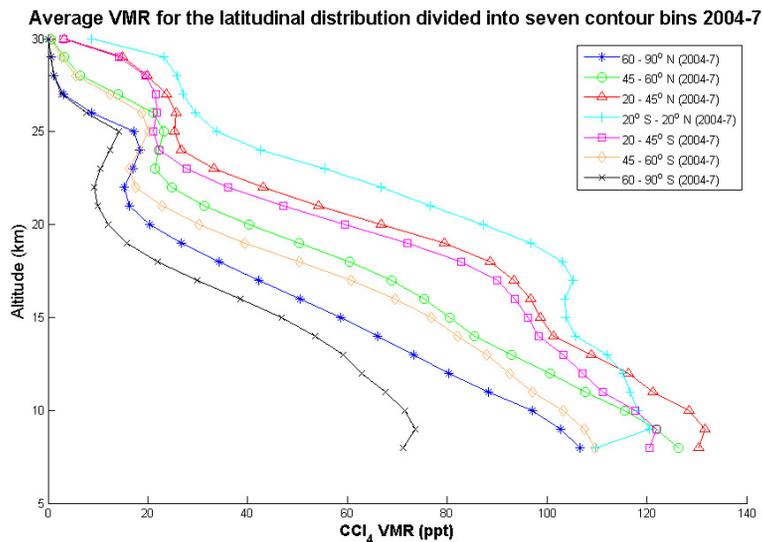
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**Fig. 2.** The locations of 8703 extravortex ACE-FTS occultations during the period February 2004 to August 2007 are indicated by coloured markers. Each colour denotes a different year; black: 2004, blue: 2005, red: 2006 and green: 2007.



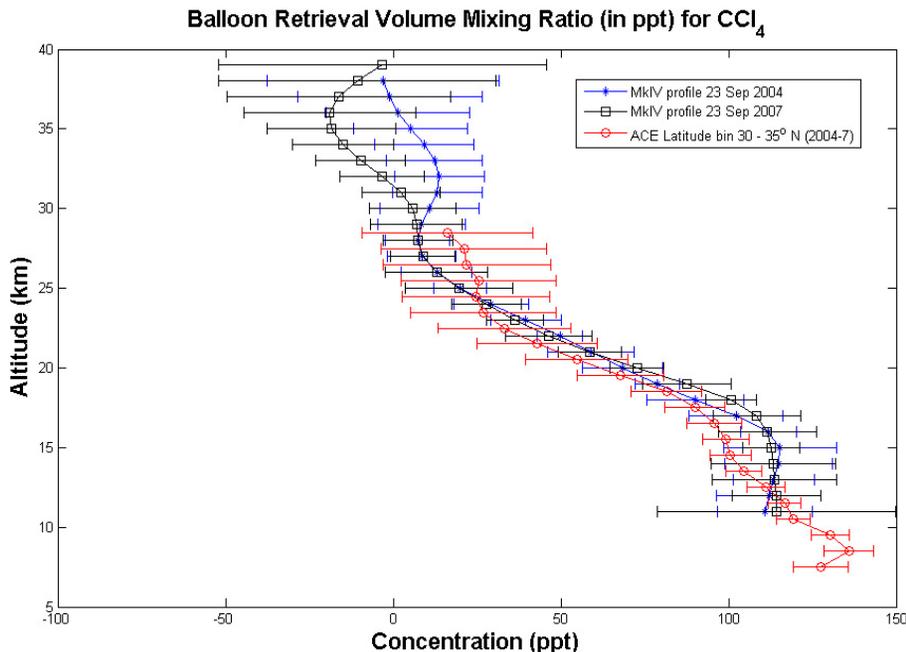
**Fig. 3.** A near global distribution of the average volume mixing ratio of carbon tetrachloride (in ppt) as a function of latitude between 8 and 30 km altitude obtained from ACE-FTS profiles from February 2004 and August 2007 is depicted. The average tropopause height is marked with a dashed black line.

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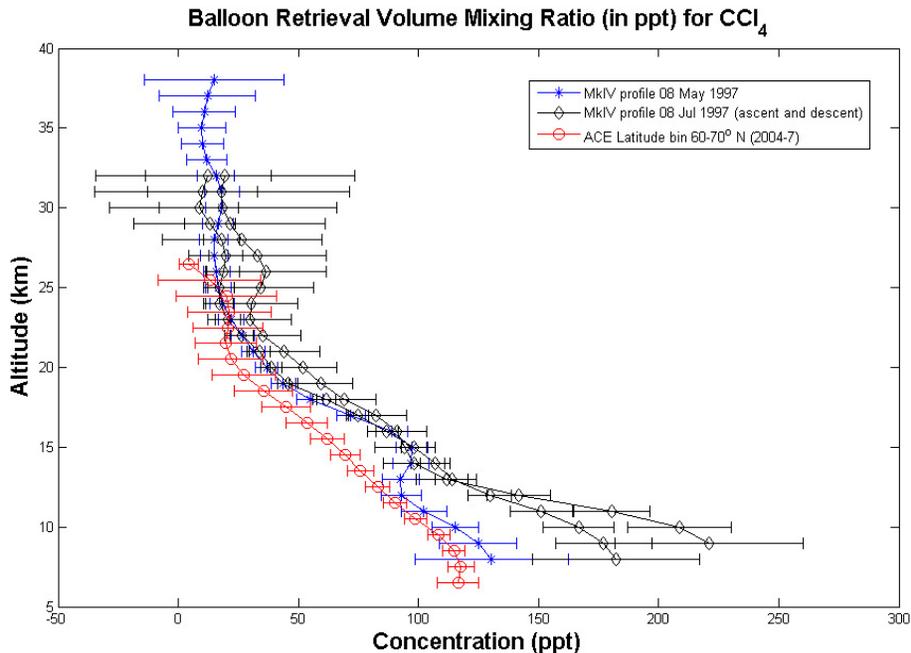
**Fig. 4.** The average volume mixing ratio (in ppt) from ACE-FTS profiles spanning February 2004 to August 2007 for seven zonal regions: 60–90° S (black line with crosses), 45–60° S (orange line with diamonds), 20–45° S (magenta line with squares), 20° S–20° N (cyan line with plus signs), 20–45° N (red line with triangles), 45–60° N (green line with circles) and 60–90° N (blue line with asterisks) are compared and displayed with respect to altitude between 8 to 30 km. Polar measurements were not retrieved above 25 km whilst in the tropics retrievals were made up to 30 km. A priori profiles have been spliced on above the maximum retrieved value so all curves reach 30 km; high altitude peaks near 25 km are artifacts.

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**Fig. 5.** The average carbon tetrachloride volume mixing ratio profile for the latitude bin 30–35° N (red line with circles) from the ACE mission between February 2004 and August 2007 is displayed and compared with two balloon profiles from near Ft. Sumner, New Mexico (33–35° N) obtained by Toon et al. (unpublished data) using a MkIV FTS instrument on the 23 September 2004 (blue line with asterisks) and 23 September 2007 (black line with squares).

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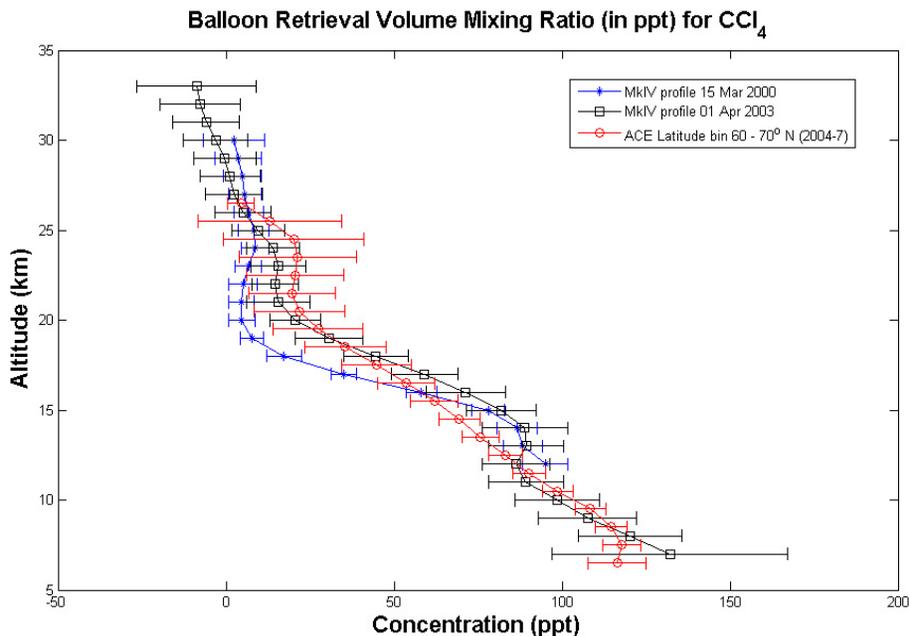


**Fig. 6.** The average carbon tetrachloride volume mixing ratio profile for the latitude bin covering 60–70° N (red line with circles) from the ACE mission between February 2004 and August 2007. The average profile is compared with three balloon profiles obtained using the MkIV FTS instrument launched from Fairbanks, Alaska (64–70° N) on the 8 May 1997 (blue line with asterisks) and 8 July 1997 (ascent and descent both black lines with diamonds) obtained by Toon et al. (1999).

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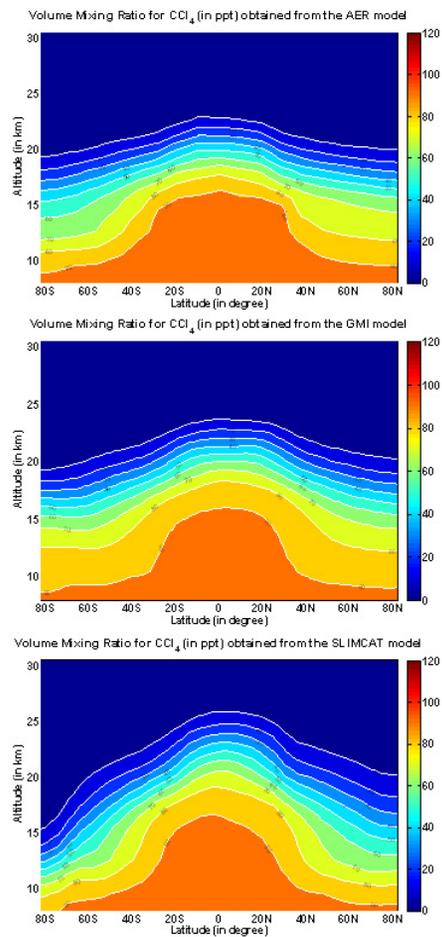
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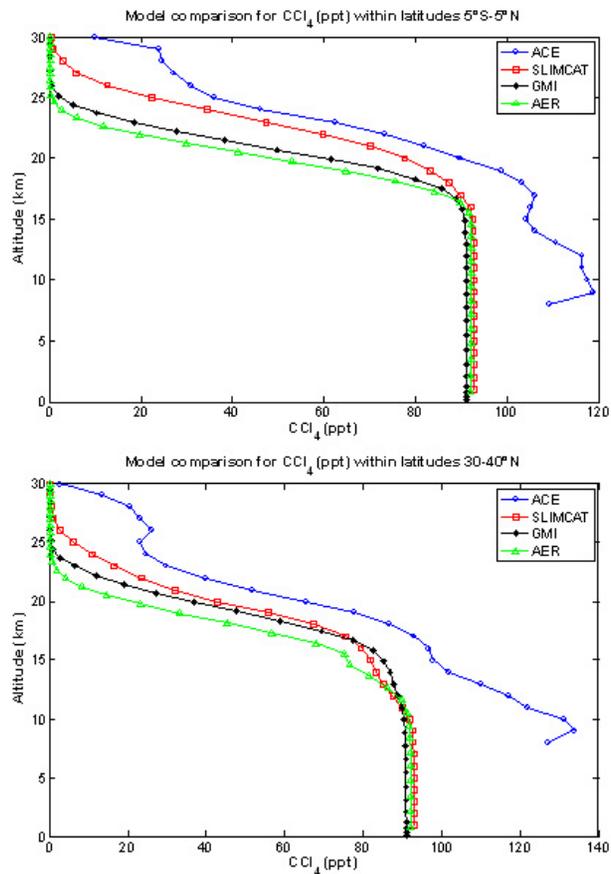
**Fig. 7.** The average carbon tetrachloride volume mixing ratio profile for the latitude bin covering 60–70° N (red line with circles) from the ACE mission between February 2004 and August 2007 is displayed. This average profile is compared with two balloon profiles from near Esrange, Sweden (62–70° N), obtained by Toon et al. (unpublished data) using the MkIV FTS instrument on the 15 March 2000 (blue line with asterisks) and 1 April 2003 (black line with squares).

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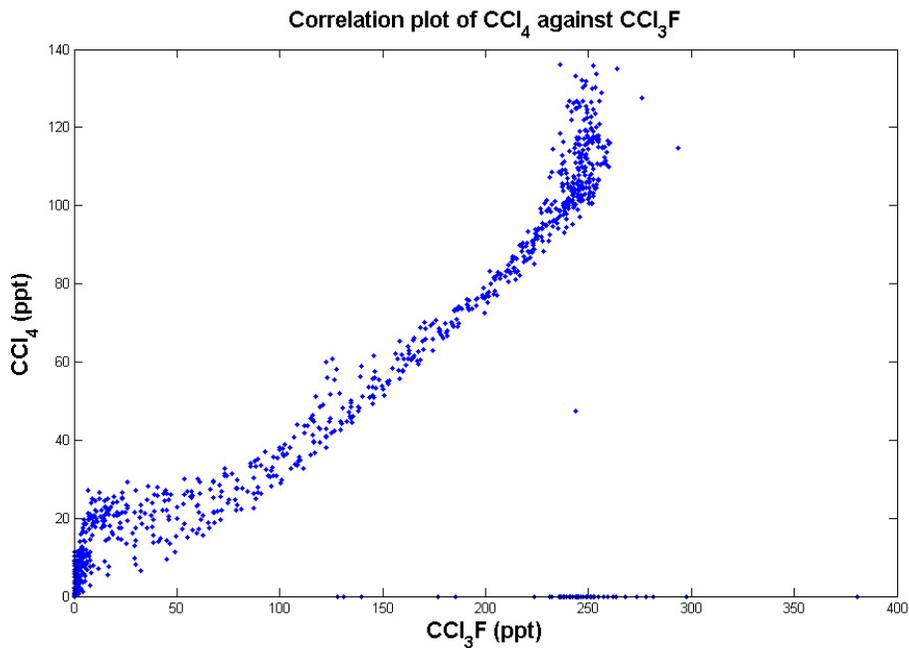
**Fig. 8.** The zonal mean annual mean latitude-height distribution of carbon tetrachloride VMR (ppt) from (a) the AER 2-D model, (b) the GMI 3-D model and (c) the SLIMCAT 3-D model.

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**Fig. 9.** The average mixing ratio of carbon tetrachloride as a function of altitude for the AER (green line with triangles), GMI (black line with asterisks) and SLIMCAT (red line with squares) models compared with the average ACE profile (blue line with circles) between the latitudes **(a)**  $5^\circ\text{S}$ – $5^\circ\text{N}$  and **(b)**  $30$ – $40^\circ\text{N}$ .

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**Fig. 10.** Correlation plot of  $\text{CCl}_4$  VMRs with  $\text{CCl}_3\text{F}$  VMRs. The lifetime of carbon tetrachloride was estimated to be  $34 \pm 5$  years from the linear part of the correlation with CFC-11.

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