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**Shortwave radiative
forcing bias of liquid
and ice clouds**

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The shortwave radiative forcing bias of liquid and ice clouds from MODIS observations

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Abstract

We present an assessment of the plane-parallel bias of the shortwave cloud radiative forcing SWCRF of liquid and ice clouds at 1 deg scales using global MODIS (Terra and Aqua) cloud optical property retrievals for four months of 2005 representative of the meteorological seasons. The (negative) bias is estimated as the difference of SWCRF calculated using the Plane-Parallel Homogeneous (PPH) approximation and the Independent Column Approximation (ICA). PPH calculations require MODIS-derived grid-point means while ICA calculations require distributions of cloud optical thickness and effective radius as well as ancillary surface albedo and atmospheric information consistent with the MODIS retrievals. With the aid of broadband solar radiative transfer algorithm we find that the absolute value of global SWCRF bias of liquid clouds at the top of the atmosphere is about 6 W m^{-2} for MODIS overpass times while the SWCRF bias for ice clouds is smaller in absolute terms by about 0.7 W m^{-2} , but with stronger spatial variability. If effective radius variability is neglected and only optical thickness horizontal variations are accounted for, the absolute SWCRF biases increase by about $0.3\text{--}0.4 \text{ W m}^{-2}$ on average. Marine clouds of both phases exhibit greater (more negative) SWCRF biases than continental clouds. Finally, morning (Terra)–afternoon (Aqua) differences in SWCRF bias are much more pronounced for ice than liquid clouds, up to about 15% (Aqua producing stronger negative bias) on global scales, with virtually all contribution to the difference coming from land areas. The substantial magnitude of the SWCRF bias, which for clouds of both phases is collectively about 4 W m^{-2} for diurnal averages, should be a strong motivation for global climate modelers to accelerate efforts linking cloud schemes capable of subgrid condensate variability with appropriate radiative transfer schemes.

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1 Introduction

In a recent study Oreopoulos et al. (2007) examined the albedo bias due to use of MODIS (Moderate Resolution Imaging Spectroradiometer)-inferred liquid cloud optical thickness and effective radius 1° gridpoint means instead of resolving their variability. This so called Plane-Parallel Homogeneous (PPH) bias (Cahalan et al., 1994a) assumed values close to 0.03, i.e., about 10% of the liquid cloud albedo with spatial cloud variations included. The motivation behind that study was the lack of an extensive global mapping of the PPH bias, with previous published satellite studies being mainly assessments of the visible PPH bias over limited parts of the globe (Barker, 1996; Oreopoulos and Davies, 1998; Pincus et al., 1999), or focusing primarily on parameters quantifying the underlying cloud horizontal inhomogeneity (Rossow, 2002; Oreopoulos and Cahalan, 2005) instead of the broadband bias itself. The current study seeks to expand the Oreopoulos et al. (2007) study by providing better seasonal coverage (using one representative month for each season instead of only winter-summer coverage) and by also including clouds classified by MODIS to be of ice phase (near their top). A newer, improved version of MODIS cloud data is used, and emphasis is placed on the shortwave (SW) Cloud Radiative Forcing (SWCRF) bias which takes into account the areal coverage and frequency of occurrence of liquid and ice clouds and directly relates the magnitude of the bias to the radiative energy budget. The SWCRF bias features presented here along with the online collection of PPH albedo biases from ISCCP (International Satellite Cloud Climatology Project) at <http://isccp.giss.nasa.gov> (from larger reference areas and assuming a different cloud classification of low, middle and high clouds) provides a fairly comprehensive picture of the radiative effects of neglecting horizontal cloud inhomogeneity. Global models that aspire to produce clouds with subgrid variability, and super-parameterization approaches (Khairoutdinov et al., 2005) should find value in these datasets.

Computational and dataset information are provided in the next section; the various dependencies of the global and local SWCRF biases are detailed in the five subsec-

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tions of Sect. 3, while conclusions, along with suggestions on how to exploit the results for global model validation, are provided in Sect. 4.

2 Dataset and radiative transfer calculations

As in Oreopoulos et al. (2007), we use daily MODIS Level-3 (1° resolution gridded) daytime data from both the Terra ($\sim 10:30$ local time overpass) and Aqua ($\sim 13:30$ overpass) satellites (datasets MOD08_D3 and MYD08_D3, respectively). Here, we use products from the most recent processing stream, Collection 5, and extend the study to four months, January, April, July, and October 2005. We extract the mean daily values of column-integrated optical thickness ($\bar{\tau}$), effective radius (\bar{r}_e), cloud fraction of successful cloud retrievals (A_c), and solar zenith angle (SZA), as well as one-dimensional (1-D) histograms of τ and joint (2-D) histograms of τ - r_e and τ -cloud top temperature (T_c), constructed by sampling every 5th pixel of the original 1 km resolution retrieval (King et al., 2003). The 1-D histograms of τ are resolved in 45 bins for liquid clouds and 30 bins for ice clouds. The 2-D histograms of τ and r_e are resolved in 110 bins (11 for τ covering the range 0.1–100 and 10 for r_e covering the range 3 to 30 μm) for liquid clouds, and 143 bins for ice clouds (11 for τ covering the range 0.1–100 and 13 for r_e covering the range 5 to 90 μm); the joint histograms of τ and T_c are resolved in 143 bins (11 for τ and 13 for T_c) for both phases. Except for high latitudes where gridpoints can be revisited within the same day due to orbital swath overlap, the daily histograms represent mainly instantaneous spatial variability of τ and r_e within the 1° gridpoint for the largest portion of the globe. The decreasing number of instantaneously viewed pixels within 1° gridpoints at high latitudes is balanced by multi-sampling due to revisits within the same day, making thus the daily cloud variability spatiotemporal instead of purely spatial. The consequences of the constant 1° grid of the MODIS Level-3 dataset and the nature of MODIS sampling in comparisons with global models is discussed later.

The radiative transfer calculations yielding daily atmospheric column albedo are performed with a version of the broadband (BB) CLIRAD-SW radiation model of Chou et

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al. (1998). The salient features of this code and the manner in which it is interfaced with the MODIS retrievals, Global Data Assimilation System (GDAS) atmospheric information (Derber et al., 1991) and MODIS-derived surface albedo (Moody et al., 2005) is described in Oreopoulos et al. (2007) and Oreopoulos and Platnick (2008). In our radiative transfer calculations, the cloud is placed in the layer whose top temperature is closest to the mean cloud top temperature (\bar{T}_c) as derived from the joint histogram of τ and T_c . The mean cloud top temperature was derived from joint histograms because the relevant Level-3 mean cloud product does not distinguish between liquid and ice phases. Since the MODIS-inferred cloud properties are assigned to a single layer of our atmospheric profile, there is no need to deal with cloud overlap which is in any case not resolved by the passive MODIS observations.

An important modification in our version of CLIRAD-SW is the introduction of a new method of calculating cloud optical properties (extinction, single-scattering albedo, asymmetry factor). The changes implemented for liquid clouds are described in Oreopoulos et al. (2007) and Oreopoulos and Platnick (2008). One of the reasons the original parameterization was changed was to extend its applicable range above the upper limit $20 \mu\text{m}$ for which it was designed originally, given that MODIS liquid effective radius retrievals can be as high as $30 \mu\text{m}$. The retrieved τ from MODIS was used in the ultraviolet-visible (UV-VIS) spectral region of CLIRAD-SW which assumes a constant cloud extinction coefficient. The spectral values of τ for the remaining three spectral regions in the SW infrared and near infrared (also with constant extinction coefficients) were found by rescaling the MODIS τ with the ratio of the extinction coefficient for those bands to its counterpart in the UV-VIS band for the appropriate retrieved value of r_e . For ice clouds a new parameterization of scattering properties is used, based on the ice particle single-scattering properties of Yang et al. (2000, 2005). The ice habit distribution is consistent with that used for the MODIS retrieval look-up tables of Collection 5, which come from Baum et al. (2005). The particle size distributions of ice clouds come from several campaigns (see details in Baum et al., 2005) and from 21 of the 30 distributions in Fu (1996). The cloud mass extinction coefficient (β), single scattering

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co-albedo ($1-\bar{\omega}$) and asymmetry factor (g) are fitted as a function of the effective ice crystal diameter D_e of the particle size distribution as follows:

$$\beta = a_0 + \frac{a_1}{D_e} + \frac{a_2}{D_e^2} \quad (1a)$$

$$1 - \bar{\omega} = b_0 + b_1 D_e + b_2 D_e^2 + b_3 D_e^3 + b_4 D_e^4 + b_5 D_e^5 \quad (1b)$$

$$g = \begin{cases} c_0 + c_1 D_e & \text{for } D_e \leq 40 \mu\text{m} \text{ and } D_e \geq 200 \mu\text{m} \\ c_0 + c_1 D_e + c_2 D_e^2 + c_3 D_e^3 + c_4 D_e^4 + c_5 D_e^5 & \text{for } 40 \mu\text{m} < D_e < 200 \mu\text{m} \end{cases} \quad (1c)$$

where the fitting coefficients a_0, a_1, a_2, b_0, b_1 , etc., are found from regressions. There is one set of coefficients for each SW infrared and near infrared spectral regions (i.e., constant single-scattering properties are again assumed within these spectral regions) and 8 sets of coefficients for the UV-VIS spectral region, one for each of its 8 bands. Thus, there are 11 sets of fitting coefficients in total. The ice optical thickness retrieved by MODIS was assigned to band 8 which covers the visible portion of the model's UV-VIS spectral region. The optical thicknesses in the remaining model bands and spectral regions were found using the same rescaling procedure described above for liquid clouds.

Similar to Oreopoulos et al. (2007), three different albedos (R) are calculated with the SW code: (1) albedos using the $\bar{\tau}$ and \bar{r}_e values of the gridbox (the PPH albedo R_{PPH}); (2) albedos using the 1-D histogram of τ and the gridbox mean value of effective radius \bar{r}_e (type 1 ICA albedo R_{ICA1}), i.e., obtained from multiple albedo calculations weighted by the relative frequency in each τ bin; and (3) albedos using the 2-D histogram (type 2 ICA albedo R_{ICA2}), i.e., obtained from multiple albedo calculations weighted by the relative frequency in each (τ, r_e) bin. The albedo calculated from the first method minus that calculated from the second gives the classic plane-parallel albedo bias with constant microphysics ($B_1^R > 0$). The albedo calculated from the first method minus that calculated from the third gives the albedo bias due to joint horizontal variations of τ and

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r_e ($B_2^R > 0$). Mathematically, the biases can be expressed as follows:

$$B_1^R(\bar{\tau}, \bar{r}_e, \nu_\tau, \mu_0) = R_{\text{PPH}} - R_{\text{ICA1}} \equiv \mathbf{R}(\bar{\tau}, \bar{r}_e, \mu_0) - \int \mathbf{R}(\tau, \bar{r}_e, \mu_0) \rho(\tau) d\tau \quad (2a)$$

$$B_2^R(\bar{\tau}, \bar{r}_e, \nu_{\tau, r_e}, \mu_0) = R_{\text{PPH}} - R_{\text{ICA2}} \equiv \mathbf{R}(\bar{\tau}, \bar{r}_e, \mu_0) - \iint \mathbf{R}(\tau, r_e, \mu_0) \rho(\tau, r_e) d\tau dr_e \quad (2b)$$

where μ_0 is the cosine of the solar zenith angle, ν is a measure of either τ or joint $\tau - r_e$ variability (e.g., a shape parameter of the 1-D probability density function $\rho(\tau)$ or the 2-D probability density function $\rho(\tau, r_e)$), and \mathbf{R} is the reflectance function (for example, the analytical solution of the two-stream approximation). The dependencies of the albedo bias on molecular absorption, Rayleigh scattering, and surface albedo are not explicitly shown in the above equations, so Eqs. (2a) and (2b) strictly correspond to the case of isolated clouds only. It should be understood, however, that all these factors (assumed to be homogeneous within the 1° gridbox) are accounted for in our calculations. Note that the ICA calculations are subject to errors due to discretization in the 1-D and 2-D histograms, but these errors are of random nature. Still, they may result in occasional negative values of bias at small values of τ where \mathbf{R} depends quasi-linearly on τ , and these are set back to zero when they occur. Since ICA albedos are based on 1-D radiative transfer calculations, they also suffer, of course, from errors due to neglect of real-world horizontal photon transfer (e.g., Cahalan et al., 1994b).

Oreopoulos et al. (2007) has shown that the albedo bias and the bias in the top-of-the-atmosphere (TOA) shortwave cloud radiative forcing $\Delta\text{SWCRF}^{\text{TOA}}$ are simply related via:

$$\Delta\text{SWCRF}^{\text{TOA}}(< 0) \equiv \text{SWCRF}_{\text{PPH}}^{\text{TOA}} - \text{SWCRF}_{\text{ICA}}^{\text{TOA}} = -A_c B^R \mu_0 S_0 \quad (3)$$

where S_0 is the incident solar irradiance at TOA. $\text{SWCRF}^{\text{TOA}}$ is defined simply as the difference in reflected solar fluxes between cloudless and all-sky (i.e., including clear-cloud mixtures) conditions (Ramanathan et al., 1989). Note that since $B^R > 0$,

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$\Delta\text{SWCRF}^{\text{TOA}}$ is a negative quantity, and does not require separate estimates of the individual PPH and ICA SWCRFs because the cloudless sky fluxes are identical and cancel out. With all forcing calculations referring to TOA in this paper, the superscript will be dropped henceforth for simplicity. Moreover, the negative sign of the SWCRF bias is dropped as well, and all magnitude comparisons are discussed in terms of absolute values.

The SWCRF bias estimates are performed for each day of the month in each grid-box where illumination conditions allow MODIS cloud property retrievals, and are then arithmetically averaged to monthly values (the impact of some averaging choices is examined later). Zonal and global averages of the gridpoint monthly values are trivially estimated as in Oreopoulos et al. (2007), but with gridpoints not receiving solar illumination contributing zero to the averages. Except for Sect. 3.5 where we explicitly examine Terra-Aqua differences, all other results correspond to averages from the two satellites.

3 Results

3.1 Overpass vs. daily and diurnal SWCRF bias

The simplest calculation of ΔSWCRF for a particular day involves combining the Terra and Aqua PPH albedo bias B^R with the insolation corresponding to the gridpoint mean SZA for that day as extracted from the MOD08_D3 and MYD08_D3 files. This SZA for most gridpoints corresponds to the SZA of the only daylight overpass for that day and is thus ~ 90 min removed from the SZA at local noon. We call the SWCRF bias obtained this way “overpass” ΔSWCRF . Since it corresponds to relatively high sun conditions it is not a good representative of the true energy impact of the neglect of horizontal cloud inhomogeneity for the duration of the entire day (sunrise to sunset). An accurate true diurnal estimate of ΔSWCRF is on the other hand not possible since the diurnal variation of cloud properties (cloud fraction and cloud properties that determine their albedo) is

not properly resolved with only two available measurements during daytime. To be able to assess, however, even crudely the influence of variable solar illumination throughout the day, we adopt the methodology of Oreopoulos et al. (2007) for calculating “daytime” Δ SWCRF’s, i.e., pairing the instantaneous PPH albedo with the instantaneous insolation at 2-h intervals, and integrating over the points in time when the sun is above the horizon. For the time period between sunrise and noon the Terra cloud retrievals are used while from noon to sunset Aqua retrievals are used, both assumed constant within their respective daytime half. These calculations of daytime Δ SWCRF are significantly more expensive computationally than the overpass Δ SWCRF calculations, involving multiple bias calculations per day for each gridpoint.

But for the SWCRF biases to be comparable with other biases or forcings that operate uninterrupted (e.g., counterpart infrared CRF biases due to neglect of horizontal cloud condensate or cloud-top temperature variations), even the daytime Δ SWCRF’s are not proper measures of the actual energy impact of cloud inhomogeneities. Rather 24-h (“diurnal”) estimate of the SWCRF biases are needed, and those can be obtained (again, as in Oreopoulos et al., 2007) by scaling the daytime biases further by the fraction of the 24-h period that the sun is above the horizon for each gridpoint.

The mean Terra-Aqua global biases of all three types of Δ SWCRF due to the combined optical thickness and effective radius horizontal variability (i.e., used in Eq. 3) are shown in stack-bar graph form in Fig. 1. The values in parentheses indicate the ratio of global mean to standard deviation for the overpass case. The ice cloud Δ SWCRF is more spatially variable than that of liquid clouds and there is a slight but distinct tendency of greater dispersion for the vernal and autumnal months compared to the winter and summer months. Due to the seasonal changes in the geographical distribution of the SWCRF bias, the latitudinal dependence of daytime length, and the non-linear nature of the global calculation, an empirical conversion of global overpass bias to global daytime or diurnal bias is not possible: the ratio of daytime to overpass global bias ranges from \sim 0.65 to 0.78, while the ratio of diurnal to overpass global ranges from 0.32 to 0.42. These values are similar to those of Oreopoulos et al. (2007) for liquid

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clouds.

Overall, liquid clouds exhibit larger ΔSWCRF than ice clouds (~ 6.1 vs. 5.4 W m^{-2} for overpass bias), with the largest disparity in January ($>1 \text{ W m}^{-2}$ for overpass bias) and the smallest in April ($<0.25 \text{ W m}^{-2}$ for overpass bias, increasing interestingly to about 0.5 W m^{-2} for daytime bias). The seasonal variability of bias is relatively stronger for liquid than for ice clouds, especially for daytime and diurnal averages. Further analysis of liquid/ice ΔSWCRF contrasts is presented in Sect. 3.3 where cloud fraction and frequency of occurrence contributions to the bias are discussed.

3.2 SWCRF bias with and without effective radius variability

As explained in Oreopoulos et al. (2007) for liquid clouds, inclusion of r_e horizontal variability in addition to τ variability, reduces ΔSWCRF because of the negative contribution to the PPH bias stemming from the weak concavity of the albedo vs. r_e curve under constant τ . Essentially, once τ variability is specified from the combined τ - r_e MODIS retrievals, the r_e spatial variability is only generating asymmetry factor and single-scattering albedo variability. This also applies for ice clouds, of course. Figure 2 contrasts liquid and ice clouds in terms of the global ΔSWCRF reduction arising from r_e spatial variability contributions, i.e., $|A_c(B_1^R - B_2^R)\mu_0 S_0|$. The global effect of r_e spatial variability is a reduction of the combined Terra-Aqua absolute value of the overpass bias by about 0.4 W m^{-2} ($\sim 7\%$) for liquid clouds and about 0.25 W m^{-2} ($\sim 5\%$) for ice clouds. Other than this, there are no major impacts in the qualitative behavior of ΔSWCRF by neglecting r_e horizontal inhomogeneity in the calculations. For example, the ratio of global mean to standard deviation decreases only very slightly when r_e variability is neglected suggesting only minor effects in the spatial patterns of the ΔSWCRF distribution. Henceforth, all SWCRF bias results will be based on B_2^R with the understanding that particle size variability is responsible for less than 10% of its magnitude.

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3.3 Cloud fraction and frequency of occurrence contributions to the monthly SWCRF bias

Equation (3) clearly indicates that the SWCRF bias of a gridpoint depends on three factors: (1) the PPH albedo bias B^R of the cloudy portion of the gridpoint; (2) the cloud fraction A_c , and (3) the solar irradiance $\mu_0 S_0$ received by the gridpoint. Δ SWCRF divided by A_c is simply the bias of the reflected TOA flux for the cloudy portion of the gridpoint. For a given incident solar flux, this allows to examine whether high (low) Δ SWCRF's come from high (low) PPH albedo biases or high (low) cloud fractions or a combination of both. Here, we identify these "per unit cloud fraction" SWCRF biases as "no CF" biases, as in "no cloud fraction was accounted for in the calculation".

Furthermore, when calculating a gridpoint's monthly mean SWCRF bias, averaging can be performed either over the number of days when clouds of a particular phase where indeed observed at the time of the overpass or over the total number of days within the month with an overpass. The latter calculation thus assigns zero contributions to the monthly Δ SWCRF from days where no cloud of the particular phase was observed. If, for example, for a gridpoint with 25 possible opportunities for observation within a month, only 14 had liquid cloud and therefore allowed estimates of liquid SWCRF bias, a monthly value of Δ SWCRF can be obtained by dividing either by 25 or by 14, with the latter calculation reflecting the monthly SWCRF bias of liquid clouds for that gridpoint "when present". This method of not accounting for the frequency of occurrence (FO) of clouds, which obviously gives higher monthly values of Δ SWCRF, was used by Oreopoulos et al. (2007) and is identified here as the "no FO" method for calculating monthly values of SWCRF bias. Our default choice in this paper (used for the results shown so far and all the results that follow, unless specifically stated otherwise) of including in the averages the zero contributions of days without clouds of a particular phase gives a fairer estimate of monthly SWCRF biases, since the ultimate impact of a forcing (and therefore its bias) depends on the frequency at which it applies. Finally, one may also be interested in the mean SWCRF bias of the cloudy

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portion of the gridpoint only for those days when cloud was present in the gridpoint. We call this the “no CF/no FO” SWCRF bias because neither cloud fraction nor frequency of occurrence is accounted for. Such a “SWCRF bias” is more closely associated with the fundamental cloud inhomogeneity characteristics driving the plane-parallel albedo bias.

The global values of all the above types of monthly Δ SWCRF are compared in the stack bar plot of Fig. 3. The white bars are correspond to the same overpass values shown in Fig. 1. Cloud fraction and method of monthly averaging have distinctly different impact on liquid and ice clouds. For instance, if averaging is performed only over the days of the month with clouds of a particular phase present within the gridpoint, the sign of the liquid-ice Δ SWCRF difference is reversed with ice clouds now having greater biases than liquid clouds (black “CF/no FO” bars). This means that, when present, ice clouds give overall larger biases than liquid clouds, partly due to larger cloud fraction as will be explained shortly, but their overall monthly Δ SWCRF is reduced because they occur less frequently. When frequency of cloud occurrence is accounted for (averaging over all days of the month with possible observations), but the bias is normalized by the cloud fraction, i.e., when the reflected flux bias of the cloudy portion is examined, the dominance of liquid over ice SWCRF bias is restored and widened (gray “no CF/FO” bars). Evidently, liquid clouds occur more frequently (at the time of the satellite overpass, strictly speaking) and are more inhomogeneous (produce large PPH cloud albedo bias, to be exact) when present. When neither days devoid of clouds of a particular phase nor the cloud fraction is accounted for in monthly estimates Δ SWCRF (striped “no CF/no FO” bars), the disparity of liquid and ice cloud tapers again because the larger frequency of occurrence of liquid clouds no longer contributes to the monthly bias. Nevertheless, the collective bias in reflected flux for the portions of the gridpoint covered by liquid clouds exceeds that covered for portions covered by ice clouds, a feature that is most prominent in January and least in July.

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3.4 Geographical distributions of the SWCRF bias

Figure 4 shows the geographical distribution of liquid and ice cloud overpass SWCRF bias from combined Terra-Aqua results for all four months. The figures reflect known patterns and regimes of the clouds of each phase and have obvious correlations with cloud inhomogeneity and PPH albedo bias maps in Oreopoulos and Cahalan (2005) and Oreopoulos et al. (2007), respectively (the latter only for liquid clouds). The largest liquid Δ SWCRF's (Fig. 4a) occur in January in the vicinity of sea of Japan and the Korean peninsula where mid-latitude winter storm systems are present, and in the eastern equatorial Pacific extending to the broader Colombia/Equador region, where cloudiness is due to shallow convection (de Szoeki et al., 2005). Neither of these two regions stands out in the other three months. The marine stratocumulus regions in the eastern parts of the major oceans also exhibit strong seasonality in Δ SWCRF, with October having in general the largest values, coinciding with the seasonal peak of cloud fraction for most of them (Oreopoulos and Davies, 1993). Mid-latitude oceans are quasi-permanent areas of large liquid SWRCF bias, but with values that are largely dependent on available solar illumination (contrast January and July southern oceans). The ice Δ SWCRF maxima on the other hand are more clearly confined in convective areas and follow the movement of the ITCZ (Fig. 4b). The mid-latitude Δ SWCRF's of ice clouds generally stay below $\sim 15 \text{ W m}^{-2}$ and are mostly smaller than their liquid counterparts, but not by as much as suggested by their color designation which is partly a result of the wider value range covered by the ice cloud colorbar.

The zonal distribution of monthly Δ SWCRF (Terra-Aqua averages) is shown only for January and July (Fig. 5). In this case we chose to show the 24-h biases to also capture latitudinal changes in sunlight duration. Features that have previously emerged in the full geographical distribution are prominent, such as the summer peaks in mid-latitude liquid SWCRF bias which assume values close to 7 W m^{-2} . The ice Δ SWCRF peaks are somewhat smaller and appear in the equatorial zone, shifting with the seasonal movement of the ITCZ. Even though plentiful solar illumination is available, minima oc-

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cur in broad subtropical zones of descending portions of the Hadley cell where deserts and marine stratocumulus regions are prevalent. The hemispherical averages are also provided in the figure in order to re-emphasize the strong role of incident solar energy in modulating the SWCRF bias: summer hemispherical values are about twice as large as winter values.

Land-ocean global overpass Δ SWCRF differences are highlighted in Fig. 6. The bias is clearly greater over oceans for both cloud types and all months with the exception of July where liquid cloud biases are very similar over land and ocean. The main reason for this seems to be the dramatic decrease of Δ SWCRF over the southern midlatitude oceans (Fig. 4a), due to the lower winter illumination. Peaks of Δ SWCRF over certain land areas such as over south Asia probably play only minor role in this July near-parity of liquid cloud biases. Besides differences in cloud heterogeneity, cloud fraction, and availability of insolation, a significant contribution to the lower land SWCRF bias must come from the brighter land surfaces which tend to reduce the cloud albedo contribution to the TOA albedo, damping therefore albedo differences between homogeneous and heterogeneous clouds.

3.5 Terra vs. aqua differences

Figure 7 show the percentage differences (normalized by the combined Terra-Aqua SWCRF bias) of Terra minus Aqua overpass SWCRF biases. Differences are in general negative (Aqua bias greater in absolute value than Terra bias), but this is much more pronounced for ice clouds, suggesting a stronger diurnal cycle with an afternoon increase in cloud inhomogeneity for this type of clouds. Liquid clouds bias differences are either near-zero (April and July) or of opposite sign (January and October). To isolate the morning-afternoon differences in cloud inhomogeneity from cloud fraction and frequency of cloud occurrence effects, the “no CF/no FO” (see Sect. 3.3) overpass Δ SWCRF relative differences are also plotted. These differences are now always negative for the liquid clouds too. Clearly, cloud fraction and frequency of occurrence reduces morning-afternoon differences due to cloud inhomogeneity alone, i.e., similar

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to ice clouds, liquid clouds tend also to be more heterogeneous in the afternoon. This is consistent with the cloud inhomogeneity factor results presented by Oreopoulos and Cahalan (2005) and the PPH albedo bias results of Oreopoulos et al. (2007).

To further understand Terra and Aqua Δ SWCRF differences that may not be revealed by global means alone, we performed additional analysis on the July 2005 liquid case exhibiting near-zero Δ SWCRF difference and the January 2005 ice case exhibiting the greatest negative bias difference. Figure 8 plots frequency distributions of Terra and Aqua Δ SWCRF for these cases. It is apparent that the near parity of Terra and Aqua July liquid Δ SWCRF is not the result of cancellations from different segments of the bias distribution. The Aqua and Terra bias histograms overlap almost perfectly before starting to diverge only at the high-end tail of the distribution corresponding to rare occurrences of large Δ SWCRF values (top panel). On the other hand, for the January ice case, histogram divergence commences at higher normalized frequencies (even though the separation point is again around the 20 W m^{-2} bin as in the liquid case). Aqua forcing biases for this case are not only overall greater, but their distribution is wider as evidenced both by the shape of the histogram and the magnitude of the standard deviation of the bias distribution (given in parentheses) which is about 20% larger than that of Terra.

Finally, we also examined whether ocean-land contrasts exist in the Terra-Aqua Δ SWCRF differences. Figure 9 explains how the global differences of Fig. 7 are ultimately determined. For liquid clouds, Terra Δ SWCRF absolute values systematically exceed (fall behind) those of Aqua over ocean (over land); for ice clouds Δ SWCRF differences are negligible over ocean and quite substantial over land with the latter obviously responsible for the negative global values in Fig. 7. That this behaviour is driven almost exclusively by morning-afternoon differences in cloud inhomogeneity over land and ocean was confirmed by plotting the counterpart of Fig. 9 for “no CF/no FO” Δ SWCRF’s (not shown): oceanic differences hovered around zero while continen-

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tal differences were strongly negative (Aqua Δ SWCRF's larger in absolute magnitude), for clouds of both phases.

4 Summary and conclusions

The global plane-parallel bias of the shortwave cloud radiative forcing SWCRF (also known as the shortwave cloud radiative effect) at 1 degree scales is examined using global MODIS (Terra and Aqua) cloud optical property retrievals for four months of 2005 representative of the meteorological seasons and a broadband shortwave radiative transfer code. The absolute value of the (negative) global SWCRF bias of liquid clouds at the top of the atmosphere is $\sim 6 \text{ W m}^{-2}$ for MODIS overpass times skewed towards near solar noon conditions, while the SWCRF bias for ice clouds is smaller in absolute terms by $\sim 0.7 \text{ W m}^{-2}$, but with stronger spatial variability. A significant contributor to liquid cloud SWCRF biases being greater is the higher frequency of occurrence of liquid clouds, which in conjunction with the higher average plane-parallel albedo bias, overcompensate for the higher cloud fraction of ice clouds, when present. If effective radius variability is neglected and only optical thickness horizontal variations are accounted for, SWCRF biases increase by about $0.3\text{--}0.4 \text{ W m}^{-2}$ on average. Rough conversions of these biases to daytime and diurnal (24-h) values yield values that are $\sim 25\text{--}35\%$ and $\sim 60\text{--}70\%$ smaller, respectively. Oceanic clouds of both phases assume larger (more negative) SWCRF biases than continental clouds. Finally, morning (Terra)–afternoon (Aqua) differences in SWCRF bias are much more pronounced for ice than liquid clouds, reaching about $\sim 15\%$ (Aqua producing stronger negative bias) on global scales, with almost all contribution to the difference coming from land areas.

If one wants to distill the analysis presented here to a single representative number of the lower limit of global SWCRF bias, the diurnal “24 h” values of Fig. 1, corresponding to combined optical thickness and effective radius variability and accounting for both cloud fraction and frequency of occurrence, may be the best choice. Taking the

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arithmetic mean of the four monthly values yields a SWCRF bias of 2.4 W m^{-2} for liquid clouds and 1.8 W m^{-2} for ice clouds. Due to the nature of MODIS observations where liquid and ice clouds cover non-overlapping portions of the gridpoint, these numbers must be added. Their total of 4.2 W m^{-2} can then be considered lower bound estimate of global SWCRF bias. The lower bound characterization is probably justified mainly by the inclusion of zero contributions from cloudless and non-illuminated areas, and to a lesser extent by the omission of the relatively small fraction of clouds classified by MODIS as “mixed” and “undetermined”. Still, a more accurate assessment requires knowledge of the full diurnal variation of cloud properties, and perhaps more sophisticated treatments of atmospheric (e.g., accounting for aerosols) and surface albedo effects.

Our global SWCRF bias values, along with the more detailed breakdown of bias behaviour seen in our full suite of results should provide a valuable validation reference for global modeling approaches that are able to generate mesoscale cloud inhomogeneity, provided that some effort is extended to simulate the MODIS worldview. This would ideally entail use of some type of “MODIS simulator” where the most obvious limitations of passive radiometry retrievals are imitated. Such limitations include, but are not restricted to, presumed unobscured views of low clouds, optical thickness integration over the entire atmospheric column, and strong dependence of cloud microphysics and phase characterization to near cloud top conditions. The simulator should also include the temporal and spatial sampling strategies of MODIS. Furthermore, it is important to keep in mind that any calculations of plane-parallel albedo or forcing bias are tied to the spatial scale at which the horizontal variability of cloud properties is considered (see Oreopoulos and Davies, 1998 for dependencies on spatial scale), so that any global model–MODIS comparison should be performed on identical grids. This would require that the global model can be run at least at 1 degree resolution, something that is probably within the capabilities of most modeling groups today. A further restriction that ideally should be obeyed is that the cloud scheme capable of subgrid cloud variability explicitly or implicitly resolves cloud fields at approximate the same scale

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(~1 km) as the pixel retrievals used to build the MODIS Level-3 statistics.

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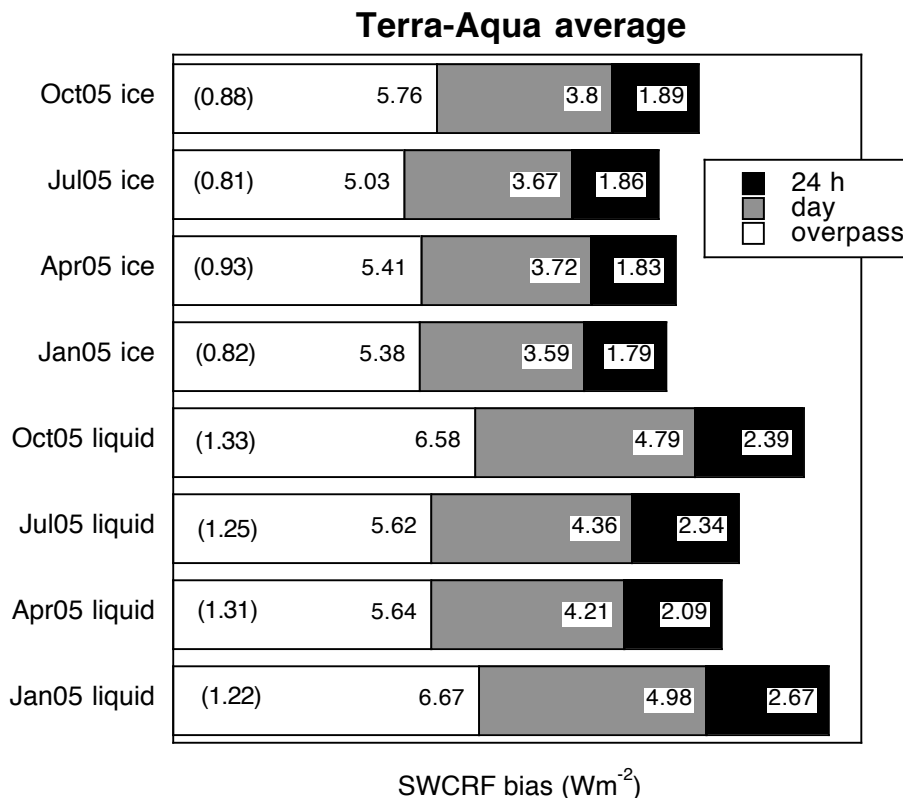


Fig. 1. Stack-bar plot showing the combined MODIS Terra-Aqua global monthly-averaged SWCRF bias using B_2^R in Eq. (3) for liquid (lower four bars) and ice (upper four bars) clouds for the four months used in this study. Overpass, daytime, and diurnal (24-h) values are shown (see text). The values in parentheses indicate the ratio of global mean to standard deviation for the overpass case.

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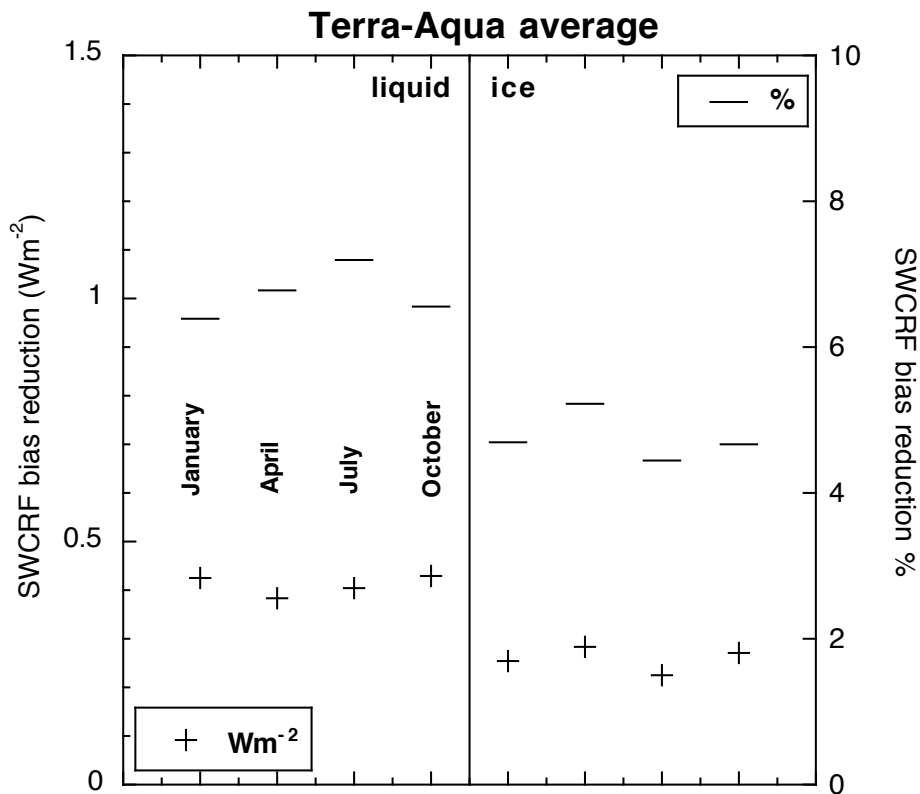


Fig. 2. Absolute ($W m^{-2}$) and percentage reduction of the combined Terra-Aqua global monthly overpass SWCRF bias from using B_2^R (r_e spatial variability included in PPH bias estimates) instead of B_1^R in Eq. (3).

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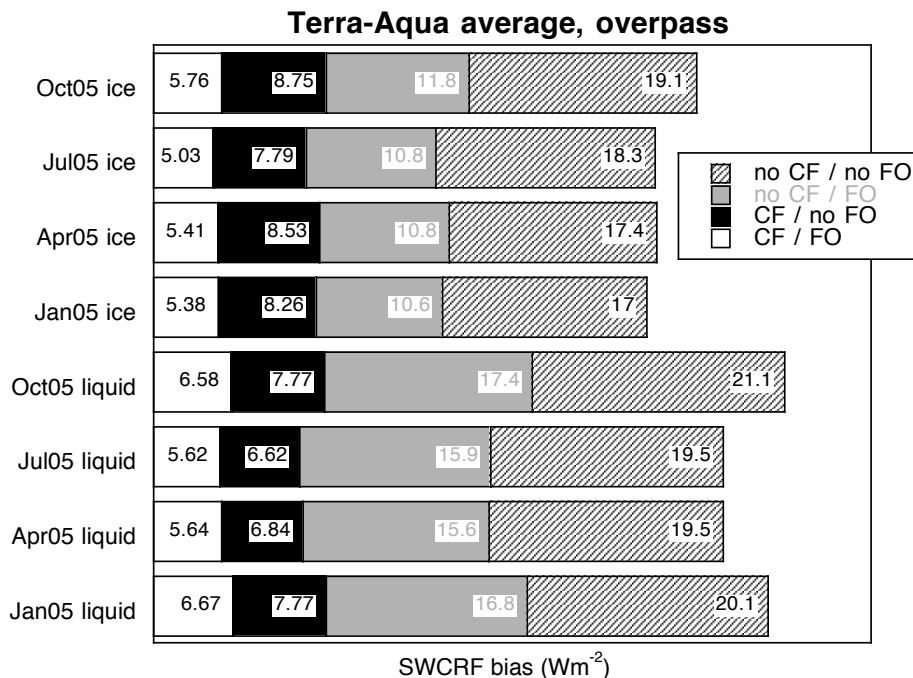


Fig. 3. Stack-bar plot showing global monthly B_2^R -based overpass SWCRF biases for our default calculation (black) and for three other methods that ignore cloud fraction and/or frequency of occurrence of clouds of the respective phase (see text for details). As in the previous plots, the values shown here are Terra-Aqua averages.

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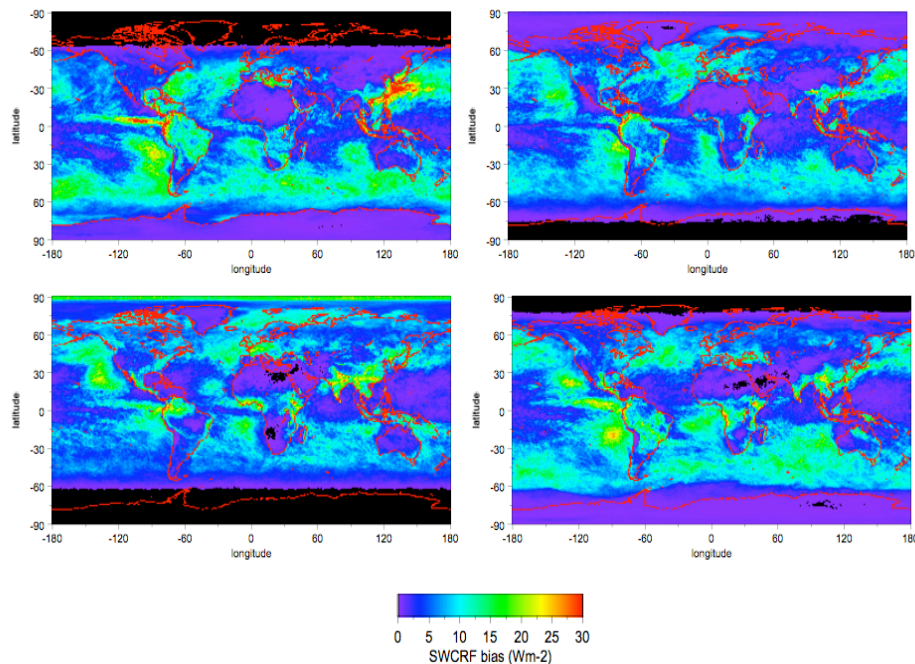


Fig. 4a. Geographical distribution of the combined Terra-Aqua monthly overpass SWCRF bias of liquid clouds from combined optical thickness and effective radius variability for the four months examined in this paper. Black areas indicate no data availability. Clockwise from top: January 2005, April 2005, October 2005, and July 2005.

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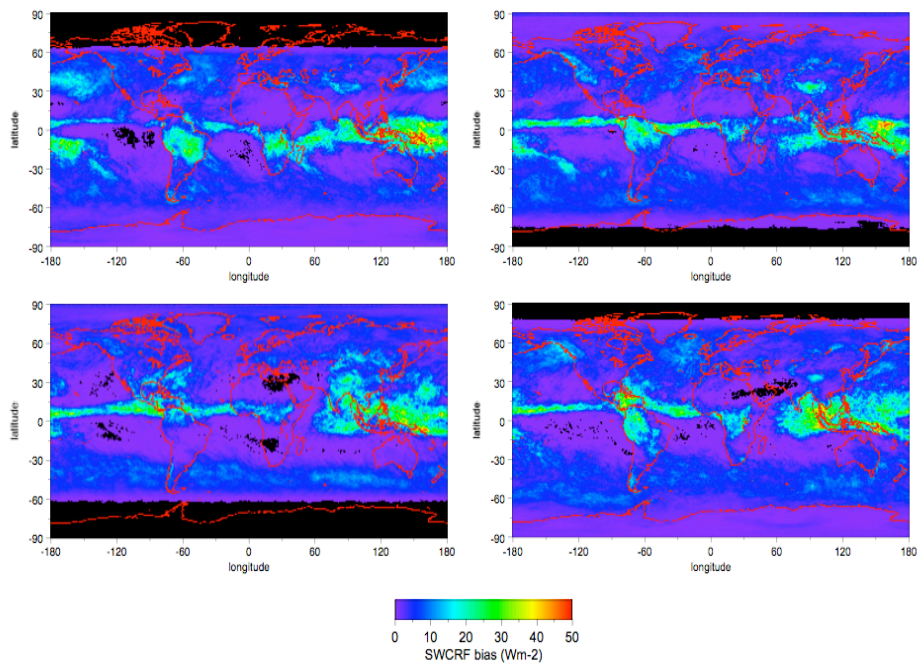


Fig. 4b. As Fig. 4a, but for ice clouds.

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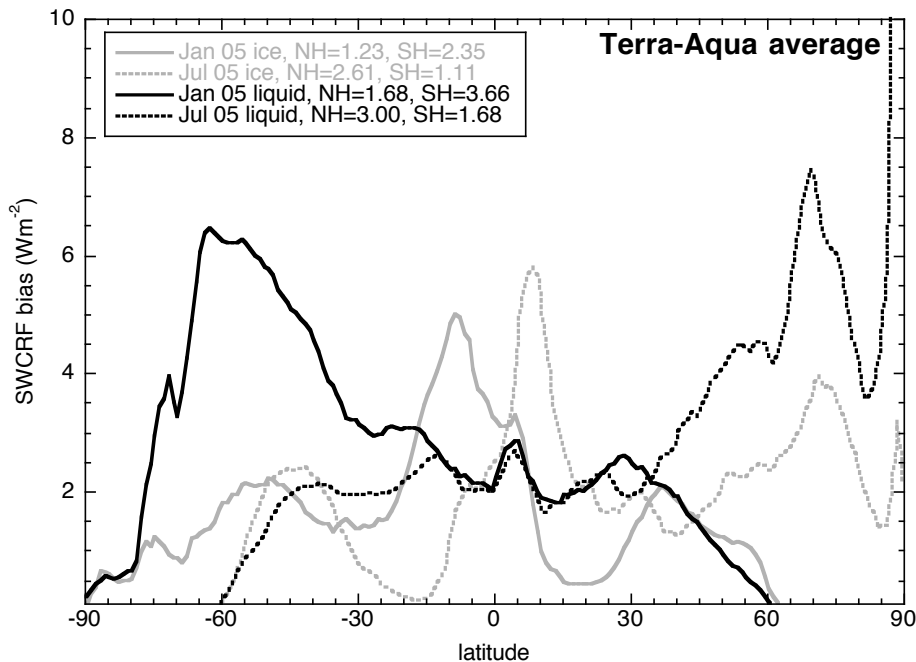


Fig. 5. Zonal dependence of the combined Terra-Aqua monthly 24-h SWCRF bias (B_2^R -based) for January and July 2005. The hemispheric biases (in $W m^{-2}$) are also provided in the legend.

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Terra-Aqua average

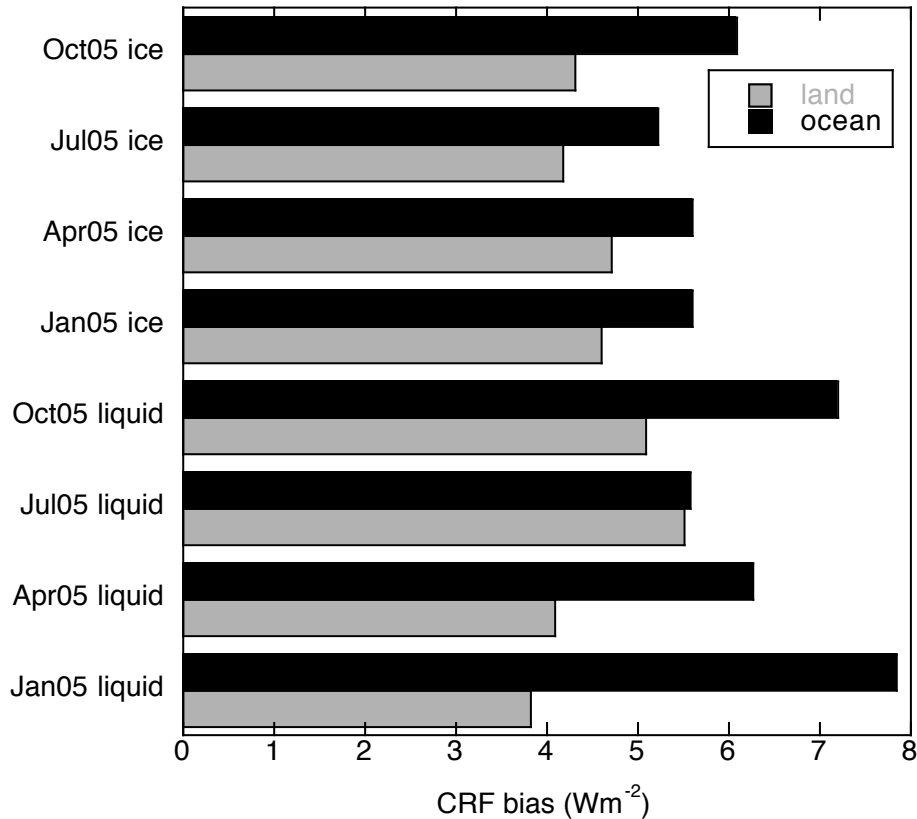


Fig. 6. Monthly combined Terra-Aqua B_2^R -based overpass SWCRF bias averaged separately over land and ocean gridpoints.

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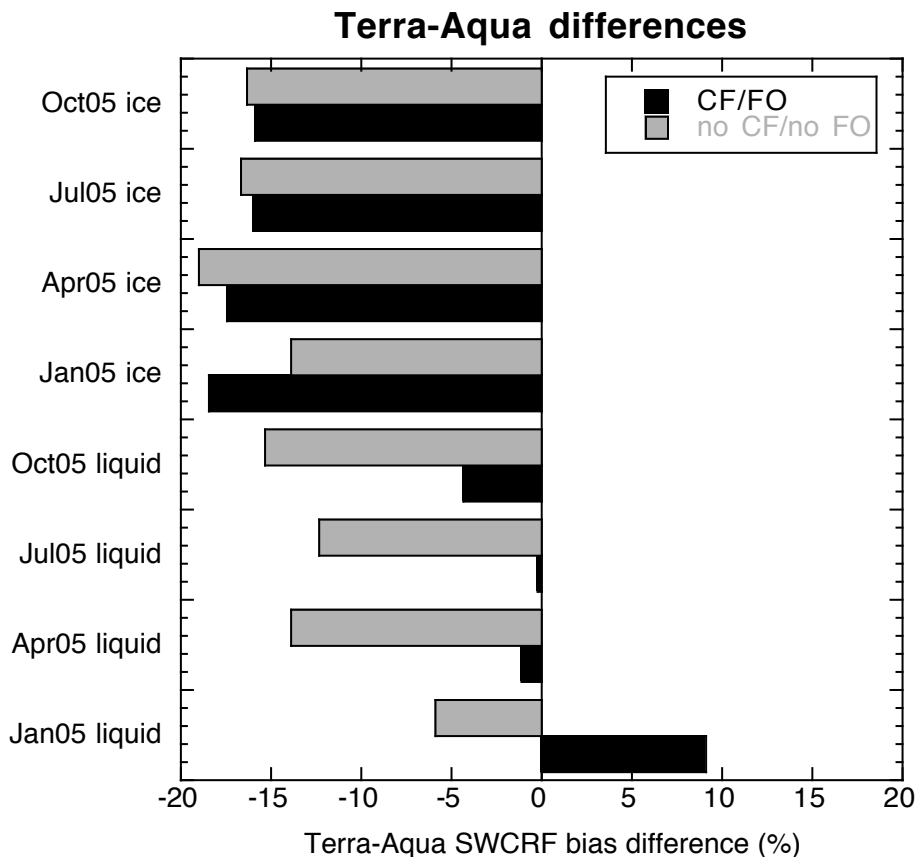


Fig. 7. Percentage difference (normalized by their combined value) of Terra minus Aqua global monthly overpass SWCRF biases (B_2^R -based). Along with the default regular SWCRF biases, results from the “no CF/no FO” (see Sect. 3.3) bias calculation are also shown. These reveal the extent to which the Terra-Aqua SWCRF biases are due to differences in cloud fraction (CF) and frequency of cloud occurrence (FO).

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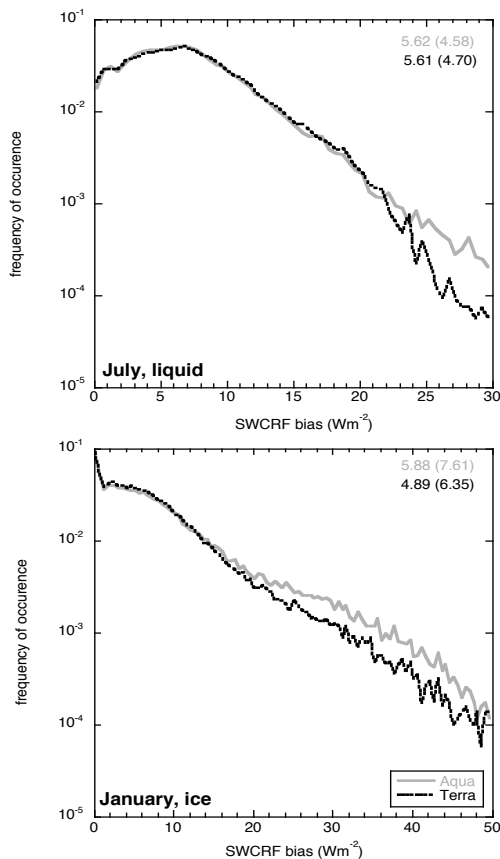


Fig. 8. Logarithmic normalized frequency of occurrence of monthly overpass SWCRF biases (B_2^R -based) for July 2005 liquid clouds (top) and January 2005 ice clouds (bottom). The global mean SWCRF biases and their standard deviations (in parentheses) are also given.

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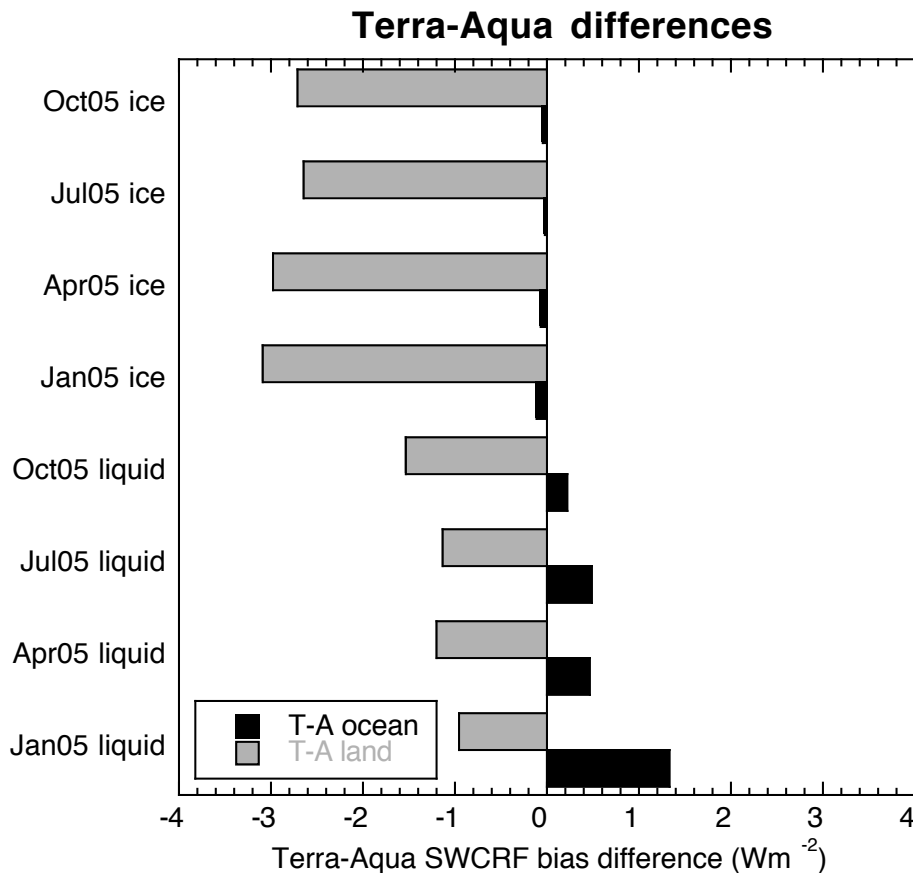


Fig. 9. Absolute difference of Terra minus Aqua monthly overpass SWCRF biases (B_2^R -based) averaged separately over ocean and land gridpoints.

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