Atmos. Chem. Phys. Discuss., 8, 6983–7016, 2008
www.atmos-chem-phys-discuss.net/8/6983/2008/

Atmospheric Chemistry and Physics Discussions

Vertical profile of peroxyacetyl nitrate (PAN) from MIPAS-STR measurements over Brazil in February 2005 and the role of PAN in the UT tropical NO_y partitioning

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Received: 18 February 2008 – Accepted: 13 March 2008 – Published: 9 April 2008

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Published by Copernicus Publications on behalf of the European Geosciences Union.

Abstract

We report on the retrieval of PAN ($CH_3C(O)OONO_2$) in the upper tropical troposphere from limb measurements by the remote-sensor MIPAS-STR on board the Russian high altitude research aircraft M55-Geophysica. The measurements were performed close to Araçatuba, Brazil, on 17 February 2005. The retrieval was made in the spectral range 775–820 cm^{-1} where PAN exhibits its strongest feature but also more than 10 species interfere. Especially trace gases such as CH_3CCl_3, CFC-113, CFC-11, and CFC-22, emitting also in spectrally broad not-resolved branches, make the processing of PAN prone to errors. Therefore, the selection of appropriate spectral windows, the separate retrieval of several interfering species and the careful handling of the water vapour profile are part of the study presented.

The retrieved profile of PAN has a maximum of about 0.14 ppbv at 10 km altitude, slightly larger than the lowest reported values (<0.1 ppbv) and much lower than the highest (0.65 ppbv).

Besides the NO_y constituents measured by MIPAS-STR (HNO_3, $ClONO_2$, PAN), the situ instruments aboard the Geophysica provide simultaneous measurements of NO, NO_2, and the sum NO_y. Comparing the sum of in-situ and remotely derived $NO+NO_2+HNO_3+ClONO_2+PAN$ with total NO_y a deficit of 30–40% (0.2–0.3 ppbv) in the troposphere remains unexplained whereas the values fit well in the stratosphere.

1 Introduction

PAN ($CH_3C(O)OONO_2$) is the most common member of peroxyacyl nitrates playing an important role in tropospheric chemistry. It is known to be eye irritant and phytotoxic to plants. PAN was firstly found in a Los Angeles photochemical smog episode (Stephens et al., 1956). Biomass burning was also suggested to be a significant source of PAN (Holzinger et al., 2005). The formation of PAN in the atmosphere involves hydrocarbons (paraffins, olefins, aromatics) and oxides of nitrogen. It is initiated by the

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reaction of OH with hydrocarbons. After intermediate reactions involving acetaldehyde, the acetyl radical and molecular oxygen, the peroxyacetyl radical (CH_3CO_3) is formed which further reacts with NO_2 to PAN (Singh, 1987).

The lifetime of PAN in the lower troposphere is in the order of hours and is dominated by thermolysis. In the upper troposphere, the lifetime, dominated by photolysis, is of the order of months, or even longer in dark Arctic regions (Talukdar et al., 1995; Kirchener et al., 1999). Median PAN/NO_y ratios of more than 0.6 at altitudes from 4 km to 8 km have been observed in long-range transported Asian pollution plumes (Roberts et al., 2004). For details on the formation and distribution of PAN see Warneck (1999) and Finlayson-Pitts and Pitts (2000).

Although PAN concentrations as high as 0.65 ppbv (up to 8 km) have been observed (Roberts et al., 2004), its typically low concentrations (<0.1 ppbv) (Tanimoto et al., 1999) make it difficult to measure. Various in situ techniques have been used to determine the volume mixing ratios of PAN in the atmosphere. These are Fourier transform infrared spectroscopy (FTIR) (Stephens et al., 1956; Hanst et al., 1982), gas chromatography with electron capture detection (GC/ECD) (Lovelock, 1961; Müller and Rudolph, 1989), gas chromatography with luminol-chemiluminescence detection (GC/LCD) (Gaffney et al., 1998), proton transfer reaction mass spectrometry (PTR-MS) (Hansel et al., 1995) coupled with a selected ion flow drift tube (SIFDT) method (Hansel and Wisthaler, 2000) and gas chromatography/negative ion chemical ionization mass spectrometry (GC/NICI MS) (Tanimoto et al., 2001).

Remote sensing in the infrared provides an alternative and independent method for the measurement of PAN. From occultation measurements of ACE-FTS on-board SCISAT-1 it was possible to retrieve PAN in a young biomass burning plume (Coheur et al., 2007). Recently, Remedios et al. (2007a) have shown the clear presence of the signatures of PAN in the emission spectra obtained by the balloon born MIPAS. Global upper tropospheric PAN distributions were derived from MIPAS/Envisat spectra by Glatthor et al. (2007).

In this paper we report the retrieval of PAN from measurements of MIPAS-STR

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(MIPAS-STRatospheric aircraft, Piesch et al., 1996) an instrument operated on board the high-altitude aircraft Geophysica. The work was initiated by the observation that large differences exist in the upper troposphere (above 10 km) between MIPAS-STR measurements of HNO_3 and coincident in situ measurements of NO_y–NO by SIOUX (StratospherIc Observation Unit for nitrogen oXides, Schmitt, 2003) also aboard the Geophysica.

In the following we give a short description of the MIPAS-STR instrument and its measurement strategy, an overview of the flight of 17 February 2005 and compare HNO_3 data from MIPAS-STR with coincident in-situ measurements of (NO_y–NO–NO_2) to obtain an upper limit PAN profile (Sect. 2). Further we give a simulation on the feasibility of detecting PAN from the MIPAS-STR observations (Sect. 3), relevant general details on the data processing (Sect. 4) and finally the retrieval of PAN from the measured spectra, including the error estimation (Sect. 5). The last section gives a summary and a discussion of the results.

2 MIPAS-STR measurements on 17 February 2005

2.1 The MIPAS-STR instrument

MIPAS-STR is a cryogenic Fourier transform emission sounder operating in the middle infrared (Fischer and Oelhaf, 1996; Keim et al., 2004). The emission method allows limb and upward viewing, yielding about 2 km vertical resolution below the flight level (up to 20 km). Reduced vertical information above the flight level is obtained by upward measurements with several elevation angles. The final results are 2-dimensional distributions of the trace gases along the flight track in an altitude range covering the lowest stratosphere and the upper troposphere.

The first deployment of MIPAS-STR was made during the Antarctic campaign APE-GAIA in 1999 (Höpfner et al., 2000). The performance of the instrument has been considerably improved in recent years.

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The pointing of the limb measurements has been operated at fixed tangent heights between 6 km and the flight altitude with a spacing of 1 km. Considering the instrumental field of view of 0.44 degrees (FWHM) over-sampling by a factor 2–3 was applied at the lower tangent heights. In addition upward measurements at elevation angles of 0, 1, 3 and 10 degrees, as well as zenith and cold blackbody (210 K) measurements were performed. Two-sided interferograms were obtained with a maximum optical path difference L of 14.4 cm, resulting in an unapodised spectral resolution ($1/2\,L$) of 0.035 cm^{-1}. For a flight altitude of 19 km the complete sequence, including calibration, takes 200 s. This results in a horizontal resolution in flight direction of about 36 km.

The data shown in this paper are obtained from channel 1, which covers the wavenumber range of 770–970 cm^{-1}.

2.2 Flight scenario

The flight track of the Geophysica with the location of the tangent points of MIPAS-STR limb sequences is given in Fig. 1. From Araçatuba (21.2° S, 50.4° W) the flight was conducted northbound and returned south on a straight track from 14° S to 23° S. Optically thick clouds were observed in the northern part of this leg which prevented trace gas retrieval for that region. However, cloudless condition were found in the southern part, just before the descent. The red rectangle in Fig. 1 indicates the tangent points of six cloud free limb sequences measured between 13:05 and 13:20 UTC. These six limb sequences cover the track of the aircraft on the descent, which started at the southernmost point of the path. The flight thus gives an excellent opportunity to compare the MIPAS-STR profile with in situ data measured during descent.

2.3 Comparison of MIPAS-STR measured HNO_3 with in situ measured NO_y–NO–NO_2

In Fig. 2 the mean HNO_3 volume mixing ratio (vmr) retrieved from the six southernmost limb sequences (see Fig. 1) is compared to the in situ observation of NO_y–NO–NO_2

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during descent. NO_y and NO are measured directly by SIOUX, while NO_2 is calculated assuming a photochemical steady state between daytime NO_2 and NO according to Eq. (1) (e.g., Schlager et al., 1997).

$$NO_2 + h\nu \rightarrow NO + O \quad \text{(R1)}$$

$$O + O_2 \rightarrow O_3 \quad \text{(R2)}$$

$$NO + O_3 \rightleftharpoons NO_2 + O_2 \quad \text{(R3)}$$

$$[NO_2] = [NO] \times [O_3] \times k(T)/J_{NO_2} \quad (1)$$

where square brackets indicate concentrations, $k(T)$ denotes the temperature-dependent rate coefficient of the reaction of O_3 with NO, and J_{NO_2} is the NO_2 photolysis frequency. The NO concentrations are taken from SIOUX measurements, the O_3 concentrations from FOZAN (Fast OZone ANalyzer, Ulanovsky et al., 2001), and the temperatures from a high-precision TDC (thermodynamic complex) sensor (Rosemount sensor customized at CAO, Central Aerological Observatory, Dolgoprudny, Russia). All three in situ instruments are aboard the Geophysica. The J_{NO_2} values are calculated with the radiative transfer model of Ruggaber et al. (1994). The vmr profile of $ClONO_2$, also included in NO_y, was retrieved from the MIPAS-STR measurements, but due to its very low vmr (see Fig. 12) neglected in the comparison.

The altitude of the cold point tropopause (see Fig. 11) is about 18 km. Below this altitude, NO_y–NO–NO_2 is always higher than HNO_3 by up to 0.32 ppbv. In the following we investigate how much of this difference can be attributed to PAN.

3 Spectral simulations for PAN

A well suited band for mid-IR PAN analysis is located between 775 and 820 cm^{-1} (Glatthor et al., 2007; Remedios et al., 2007a). To indicate the contribution of different atmospheric trace species in this spectral region we show simulations performed with KOPRA (Karlsruhe Optimised and Precise Radiative transfer Algorithm, Stiller et

al., 2000) for a tangent height of 8 km located at the southern part of the flight. In total 38 different trace gases are considered, the 14 strongest radiances are shown in Fig. 3. The band of PAN is mainly interfered by CO_2, H_2O, O_3, CCl_4, CFC-22, CFC-113, CH_3CCl_3 and $ClONO_2$. The infrared cross-section data for PAN at 295 K (Allen et al., 2005a) and the newer data at lower temperatures of 273 K and 250 K (Allen et al., 2005b) are adopted in this paper. For the simulation, the temperature profile is taken from the ECMWF model and the vmr profiles for all gases except for H_2O and PAN are taken from a tropical climatology (Remedios et al., 2007b). For PAN a midlatitude profile of the MOZART Model (Model for OZone And Related chemical Tracers, Horowitz et al., 2003) is used. The water profile is estimated from in situ measurements of FLASH (FLuorescence Airborne Stratospheric Hygrometer, Sitnikov et al., 2007) and FISH (Fast In situ Stratospheric Hygrometer, Zöger et al., 1999) aboard the Geophysica (see Fig. 5).

The sensitivity of the MIPAS-STR observation on PAN is demonstrated by plotting simulated difference spectra (with – without PAN) for various tangent heights between 8 and 18.6 km (see Fig. 4).

In small spectral regions, the information on PAN is reduced due to saturation caused by interfering trace gases. This is the case around the CO_2 Q-branch (792 cm^{-1}) and at the position of strong CO_2 and H_2O lines. Apart from these regions the radiance abates rather quickly with increasing tangent height. At 13, 14, and 15 km, the maximum radiance is only 50, 25, and 12.5 $nW/(cm^2\,sr\,cm^{-1})$, respectively, comparable with the spectral noise (15 $nW/(cm^2\,sr\,cm^{-1})$) in the single MIPAS-STR spectra.

However, with the high resolution spectra the broadly emitting PAN can be retrieved by multi-line retrieval from much lower radiances. In the present work 1171 independent spectral points were used to obtain a PAN profile.

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4 Data processing

4.1 Level-1 processing

Level-1 processing of the MIPAS-STR data provides the input data for the subsequent profile retrieval. Basically, it converts raw interferograms of the atmospheric measurements stored during the flight into radiometrically calibrated atmospheric spectra for each tangent height or elevation angle. The spectral gain and offset of the instrument were obtained from the zenith and cold blackbody measurements of each individual sequence. The zenith spectra were corrected for the contained atmospheric features. Level-1 processing also provides the auxiliary data which are derived from the stored housekeeping information as well as from the line of sight calibration and the field of view measurements made before and after the flight. The auxiliary data include information on the corrected flight altitudes, elevation- and azimuth angles, and relevant instrument parameters.

4.2 Level-2 processing

Vertical profiles of the atmospheric parameters (vmr of gases, temperature, pressure and absorption/emission of aerosols) are retrieved by use of the atmospheric radiative transfer model KOPRA and its inversion algorithm KOPRAFIT. The profiles are iteratively changed to minimise the residuum between measured spectra and forward calculated spectra of a complete sequence. Regularisation of the profile shape against an a priori profile is necessary for each retrieved atmospheric parameter because the chosen retrieval grid (0.5 km) is finer than the achievable vertical resolution.
In KOPRAFIT the Tikhonov-Philips regularisation method (Tikhonov, 1963; Phillips, 2003) was adopted:

$$\boldsymbol{x}_{i+1} = \boldsymbol{x}_i + \left(\mathbf{K}_i^T \mathbf{S}_y^{-1} \mathbf{K}_i + \gamma \mathbf{L}^T \mathbf{L}\right)^{-1} \left[\mathbf{K}_i^T \mathbf{S}_y^{-1} \left(\boldsymbol{y} - f\left(\boldsymbol{x}_i\right)\right) + \gamma \mathbf{L}^T \mathbf{L} \left(\boldsymbol{x}_a - \boldsymbol{x}_i\right)\right] \quad (2)$$

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where i denotes the iteration index; $\boldsymbol{x}$ the vector with the unknowns; x_a the a priori values; $\boldsymbol{y}$ the measurement vector; $\mathbf{S_y}$ the measurement covariance matrix of $\boldsymbol{y}$; f the forward model; $\mathbf{K}$ the spectral derivatives matrix; γ the regularisation parameter and $\mathbf{L}$ the first derivative regularisation operator.

The regularisation strength is chosen as small as possible, just to avoid oscillations in the resulting profile.

The achieved vertical resolution of the retrieved profile is the FWHM (full width at half maximum) of the columns of the averaging kernel matrix, given by:

$$\mathbf{A} = \left(\mathbf{K}_i^T \mathbf{S}_y^{-1} \mathbf{K}_i + \gamma \mathbf{L}^T \mathbf{L}\right)^{-1} \mathbf{K}_i^T \mathbf{S}_y^{-1} \mathbf{K}_i \quad (3)$$

4.3 PAN retrieval method

Here we describe the strategy used for the retrieval of PAN. To minimise the error contribution from spectral noise, we have averaged all spectra of the same tangent height/elevation angle within the six southernmost limb sequences (see Fig. 1), which reduces the noise from 15 to 6 nW/(cm^2 sr cm^{-1}). Furthermore, we have used all spectral points between 775 and 820 cm^{-1}, with the exception of the region 790–794 cm^{-1}. We excluded this interval to avoid any error on the retrieval from line-mixing of the CO_2 Q-branch located there.

A summary of atmospheric parameters (12 species and temperature) that have been considered in the retrieval scheme is given in Table 1. Among those parameters, five species (CH_3CCl_3, CFC-113, CFC-22, CFC-11, and $ClONO_2$) have been determined in steps previous to the PAN retrieval and are kept constant. The remaining profiles are fitted simultaneously with PAN.

$ClONO_2$ is fixed to the profile derived from the nearby ν_4 Q-branch in the interval 779.5–781 cm^{-1}. CFC-11 has been determined on the basis of the major band in the interval 838–856 cm^{-1} and CFC-22 has been obtained from its signature at 828.7–829.4 cm^{-1}. CH_3CCl_3 and CFC-113 profiles are firstly estimated from the tropic climatology and then scaled to remove their spectral signatures from the residuum. The

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scaling factor corrects the profiles for the annual decrease.

The vmr profiles of all five pre-determined species are plotted in Fig. 12.

As a priori vmr profile for PAN ($PAN_{a\,priori}$) a mid-latitude profile of the MOZART model is used (see Fig. 6). Beside trace gases and temperature we determine a continuum extinction profile for aerosols and a tangent height constant radiation offset for minor calibration errors.

In the retrieval we have considered spectra of tangent heights from 8 km upwards, because the lower spectra are contaminated by clouds.

4.4 Determination of the H_2O a priori profile

Although H_2O is simultaneously fitted with PAN, an impact of the applied a priori profile for water vapour on the PAN result has been observed. The use of a climatological H_2O a priori profile resulted in instabilities in the PAN vertical distribution. This was caused by the incorrect vertical position of the hygropause mapped into the resulting water vapour profile through the Tikhonow-Phillips regularisation constraint. To solve this problem we adopted a 2-step approach. In the first step we use a zero a priori H_2O profile and a relatively strong regularisation. This leads to a H_2O profile (H_2O_{first}) with reasonable position of the hygropause but relatively low vertical resolution. Its values are found to be higher than the in situ data between 10 and 12 km.

In the next step with weakened constraint, H_2O_{first} is used as the a priori to get the next H_2O profile. This profile has been used as the "selected" H_2O a priori in the PAN retrieval. As shown in Fig. 5 the fitted H_2O vmr profile is very similar to the selected a priori profile above 11 km but larger at lower altitudes. Both the fitted and the selected a priori profile tend to have some instability around 12–13 km. Such kind of feature is also present in the in situ data observed by the instruments FLASH and FISH (Fig. 5) just above the hygropause and, thus, might be real.

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4.5 PAN cross sections

Cross sections for PAN have been measured at 295 K, 273 K and 250 K (Allen et al., 2005a,b) whereas the relevant temperature for our measurements is between 197 K and 250 K. The cross sections increase from 273 K to 250 K by 8%, from 295 K to 273 K by 12%. For the PAN profile we extrapolated linearly, using the cross section measured at 273 K and 250 K, according to the atmospheric temperature profile.

4.6 The resulting PAN-profile

Figure 6 shows the retrieved profile of PAN from MIPAS-STR (in the following referred to as $PAN_{MIPAS-STR}$). Error bars indicating the total error (see Sect. 4.7 and Fig. 10) and the noise error are added in the profile. The vmr profile peaks at 10 km altitude with a value of about 0.14 ppbv and an error of 15%. At 14 km, still 0.04 ppbv of PAN are observed with an error of proximately 22%.

The averaging kernel matrix for $PAN_{MIPAS-STR}$ is used to determine the sensitivity of the retrieval at different altitudes (see Fig. 7). The columns of the matrix are the answers of the retrieval to a delta function in the associated altitude. The diagonal structures in the altitude range of the limb sequences between 8 and 18.6 km is clearly visible in Fig. 7. Below this range no measurements are available. The vertical resolution, determined as FWHM of each column of the averaging kernel matrix is given in Fig. 8. Above the flight level of 19 km the diagonal structure broadens strongly showing that there the vertical information is strongly reduced compared to the limb-range where a vertical resolution of 2–2.5 km has been achieved (see Fig. 8).

4.7 Residual spectra

We investigate the quality in the spectral domain of the PAN retrieval described in Sect. 4.3 (RUN_{fit}) in comparison with that resulting from two further approaches (see Table 2). For the test "RUN_{limit}" we fixed the PAN profile to PAN_{limit} (see Fig. 2) and

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retrieved all other parameters like described in Sect. 4.3. The test "RUN_{zero}" has been handled similarly but all PAN vmrs are fixed to zero.

For all three runs, the residual spectra are shown in the lower panels of Fig. 9 for two selected tangent heights, 11 km (left part) and 13 km (right part). The top panels show the corresponding measured spectra. The rms (root mean square) of the residuum is considerately lower [14.8 nW/(cm^2 sr cm^{-1})] for the run RUN_{fit}, than for RUN_{limit} [32.1 nW/(cm^2 sr cm^{-1})] and RUN_{zero} [27.3 nW/(cm^2 sr cm^{-1})]. The rms of RUN_{fit} is higher than the spectral noise (6 nW/(cm^2 sr cm^{-1})), because the residuum still contains residuals of lines, especially for low tangent heights. The broadband structure similar to the PAN contribution (see Figs. 3 and 4), present in the residua of RUN_{limit} and RUN_{zero}, is, however, removed in RUN_{fit}.

4.8 Error estimation

In this section we analyse the effects of various error sources on the retrieved PAN vertical profile. We distinguish in instrument-related error sources such as calibration and spectral noise and in retrieval-related like spectroscopy and the errors in the used profiles. Here we consider temperature, water vapour, CCl_4 and the five interfering species (CH_3CCl_3, CFC-113, CFC-22, CFC-11, and $ClONO_2$) whose profiles have been kept constant during the PAN retrieval. Figure 10 presents the total error together with the individual errors described in the following paragraphs.

1. Temperature: A comparison of the retrieved vertical temperature profile from MIPAS-STR with that of ECMWF and in situ observations by the Rosemount TDC is shown in Fig. 11. In general, good agreement is found between all profiles, providing us the confidence in the level-1 processing for the spectral band in which also PAN is retrieved. Since the MIPAS-STR temperature is still slightly lower in the comparison, especially in the lower part, the contribution from a 2 K shift of the temperature profile is considered in the PAN error estimation.

2. Water vapour: Two different H_2O a priori profiles ($test_1$ and $test_2$ in Fig. 5), are used to estimate the contribution of the H_2O a priori profile on the PAN error budget. In both a priori test profiles the zigzag at 13 km is removed. Additionally, the a priori values in $test_2$ have been increased for altitudes below 11 km, adapting the FISH measurement. $Test_1$ only weakly influences $PAN_{MIPAS-STR}$, whereas $test_2$ leads to differences in the order of about 5%.

3. The five pre-determined species: An uncertainty of 5% in each of the vmr profiles (CH_3CCl_3, CFC-113, CFC-22, CFC-11, and $ClONO_2$), which have been determined in previous steps and kept constant during the PAN retrieval, is assumed.

4. PAN cross sections: To consider atmospheric temperatures lower than 250 K we linearly extrapolated the cross sections measured at 273 K and 250 K. For the error from the PAN cross section, we added the temperature dependent term (T–250 K)×0.16% to the error of 3% given by Allen et al. (2005b) for 250 K. The first term, roughly 4 % for 25 K difference is the dominant term at temperatures close to 200 K.

5. Radiometric calibration: An error in the gain calibration of 2% has been assumed.

6. Spectral noise: A NESR (noise equivalent signal radiance) of 6 nW/(cm^2 sr cm^{-1}) has been assumed.

7. The a priori profile of PAN: The influence of the chosen a priori profile on the retrieved PAN has been investigated by using a zero profile instead of $PAN_{a\,priori}$.

Figure 10 presents each individual error contribution together with the total error calculated from these by the root square sum of all individual errors for each altitude. The high relative errors are in altitudes with low vmr values (see Fig. 6). In the altitude range spanned by the tangent points from 9 km to 18 km, the total relative error is between 15 % and 20 %.

In the lower part (up to about 14 km), errors in the temperature and PAN cross section dominate, whereas above spectral noise and PAN cross sections are the major error sources. Error bars for the total error are given with the $PAN_{MIPAS-STR}$ profile in Fig. 6.

5 Discussion

This work was initiated by the comparison of the MIPAS-STR HNO_3 profile with the difference profile NO_y–NO, measured by the in-situ instrument SIOUX. The disagreement between the two profiles posed the question, which of the constituents of NO_y have to be considered additionally. The profile of $ClONO_2$ was retrieved from the MIPAS-STR spectra, and NO_2 was calculated from O_3 and NO. However, the consideration of those two gases did not change the situation, as their vmrs are very small. So we tried successfully to retrieve PAN vmrs from the MIPAS-STR spectra. But the retrieved PAN profile only accounts for a sixth to a half (depending on the height) of the deficit NO_y–NO–NO_2–HNO_3–$ClONO_2$. In Fig. 13 we show all derived profiles of the individual constituents of NO_y, their sum and the measured NO_y profile.

In altitudes above the tropopause, where the tropospheric constituents of NO_y can be neglected, the vmr of NO_y–NO agrees well with the vmr of HNO_3+$ClONO_2$. This gives us confidence in the accuracy of both measurements, although the individual errors both for NO_y–NO and HNO_3+$ClONO_2$ are about 15%.

Murphy et al. (2004) report on two compounds (HO_2NO_2 and $CH_3O_2NO_2$) which become important at low temperatures in the upper troposphere. They determined their contribution to NO_y to be 30% and more at temperatures below 230 K. This could explain the discrepancy, as these compounds are measured in the sum NO_y, but can

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not be measured or calculated individually by the instruments aboard the Geophysica.

6 Conclusions

We investigated the retrieval of the vertical profile (8–19 km) of PAN using MIPAS-STR tropic emission spectra obtained in February 2005. The largest peak in the retrieved PAN vmr profile is located at 10 km altitude with an amount of about 0.14 ppbv. Above 10 km PAN decreases with a second smaller maximum at 16 km (≈0.06 ppbv). The total relative error is estimated to be about 15–20% between 9 and 18 km. Our measurements took place in February, the end of the dry season when the biomass burning was almost finished. This may explain the low values compared to measurements by (Singh, 1987, 0.3 ppbv @ 3–11 km) and (Glatthor et al., 2007, 0.33 ppbv @ 8 km and 0.23 ppbv @ 11 km) in the same region (tropic southern Atlantic) but in September/October.

Acknowledgements. Financial support for this project by the European Space Agency (contracts 10249/01/NL/SF and 16039/02/NL/SF) and the European Commission (contract EVK2-CT-2001-00122) is acknowledged. We thank NILU for providing ECMWF analysis data via the NADIR database. We thank the colleagues of FISH, FLASH, FOZAN, SIOUX, and TDC for providing their measurements.

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Table 1. Adjusted atmospheric parameter during the PAN retrieval.

Parameter	Handling	Source of a priori profile
Temperature	Cofitted	ECMWF
PAN	Cofitted	model (MOZART) ($PAN_{a\ priori}$)
H_2O	Cofitted	Pre-determined (see Sect. 4.5)
O_3	Cofitted	Climatology
CCl_4	Cofitted	Climatology
HCN	Cofitted	Climatology
C_2H_6	Cofitted	Climatology
NH_3	Cofitted	Climatology
$ClONO_2$	Pre-determined	MIPAS-STR
CFC-11	Pre-determined	MIPAS-STR
CFC-22	Pre-determined	MIPAS-STR
CH_3CCl_3	Pre-determined	Modified Climatology
CFC-113	Pre-determined	Modified Climatology

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Table 2. Three test cases for examination of the spectral fit quality.

	A priori PAN profile	Treatment of PAN in KOPRAFIT
RUN_{fit}	Model	fitted
RUN_{limit}	NO_y–NO–NO_2–HNO_3	not fitted
RUN_{zero}	Zero	not fitted

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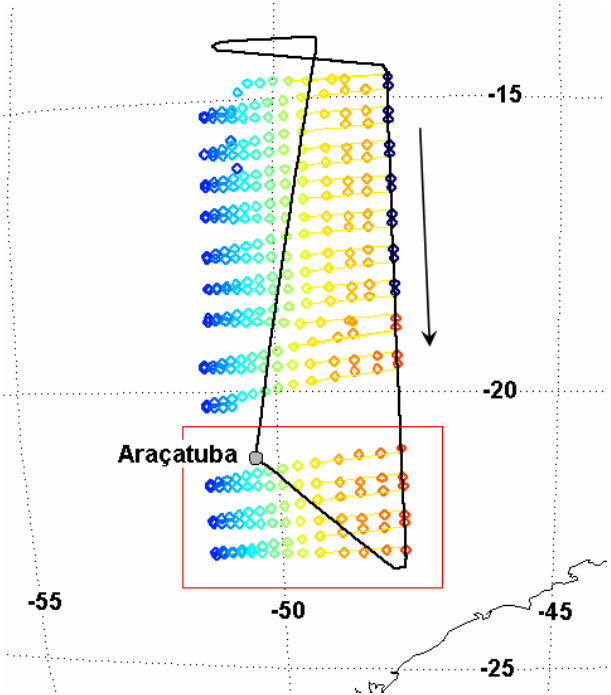

Fig. 1. Flight track of Geophysica and location of the tangent points of the MIPAS-STR instrument on 17 February 2005. The black arrow denotes the flight direction. The colour coding of the tangent points indicates their altitude, from blue for the lowest altitude at 6 km to orange at the aircraft flight level at 19 km. The red rectangle surrounds the six cloud-free sequences, used in this work.

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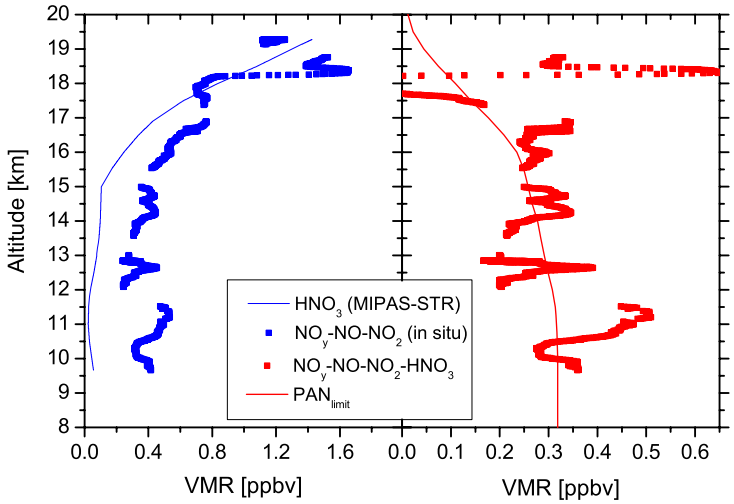

Fig. 2. Comparison of HNO_3 from MIPAS-STR with in situ measured NO_y–NO–NO_2 (left) and the difference of NO_y–NO–NO_2–HNO_3 (right). In the right panel we give also the smoothed difference used in Sect. 4.6.

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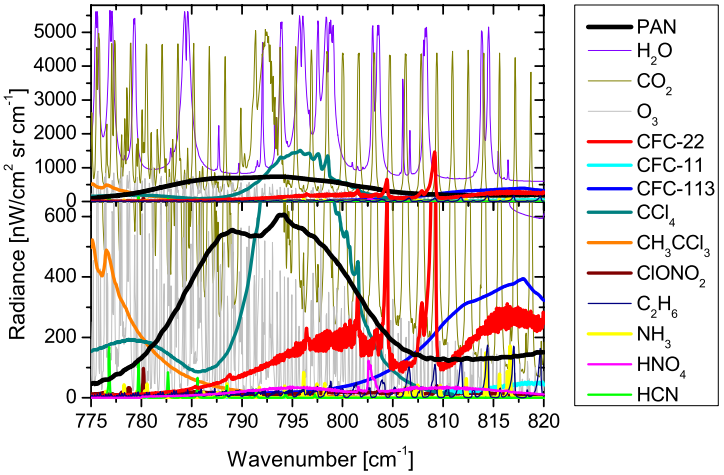

Fig. 3. Simulated spectra in the broad-band range of PAN (black) for a tangent height of 8 km. The lower plot is a zoom of the upper one in y-direction.

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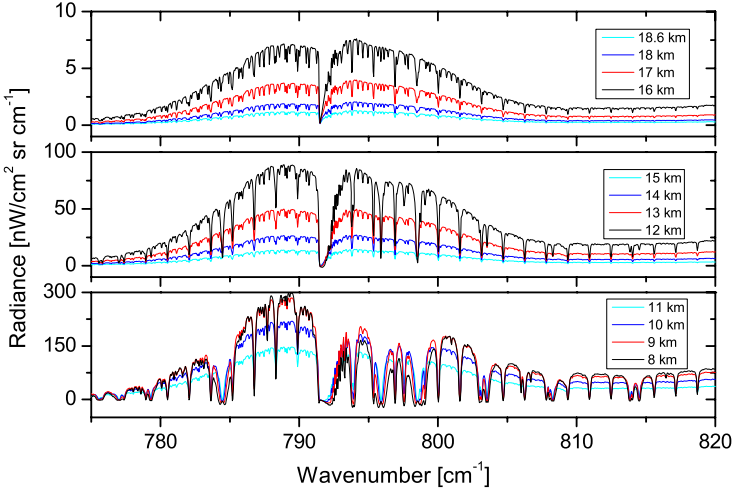

Fig. 4. Simulated difference spectra (with – without PAN) in the broad-band range of PAN for all tangent heights from 8 to 18.6 km. The noise level of MIPAS-STR is about 15 nW/(cm^2 sr cm^{-1}).

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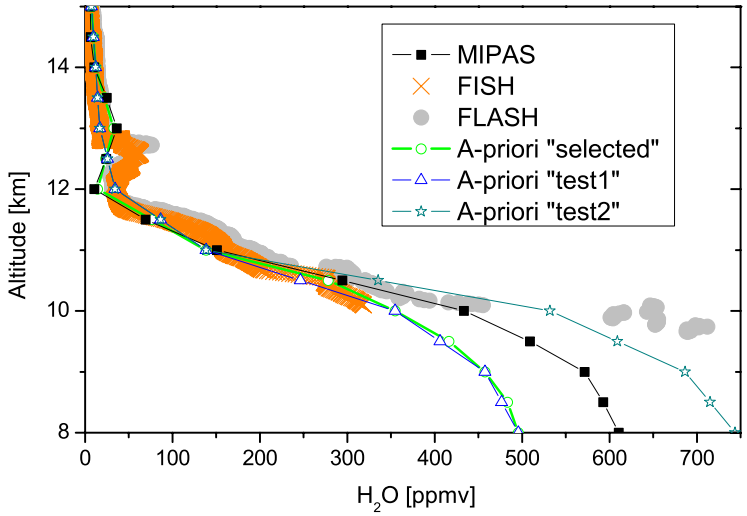

Fig. 5. The retrieved H_2O profile (simultaneously fitted with $PAN_{MIPAS-STR}$) and the selected a priori profile are shown together with in situ data from the FLASH and FISH instruments and two modified a priori profiles ($test_1$, $test_2$), used for error estimation.

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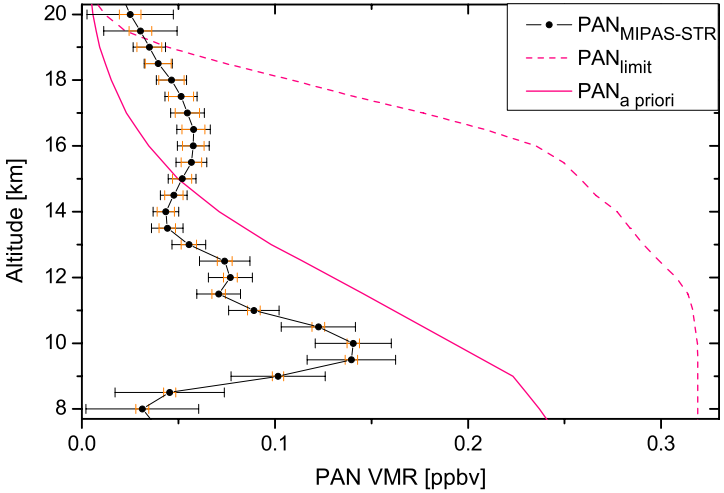

Fig. 6. PAN vertical profile retrieved from MIPAS-STR (17 February 2005, 13:15 UTC; location: 22.0° S and 47.7° W). The error bars give the noise error (red) and the estimated total error (black). Also shown are the a priori profile and the upper limit corresponding to the NO_y measurements.

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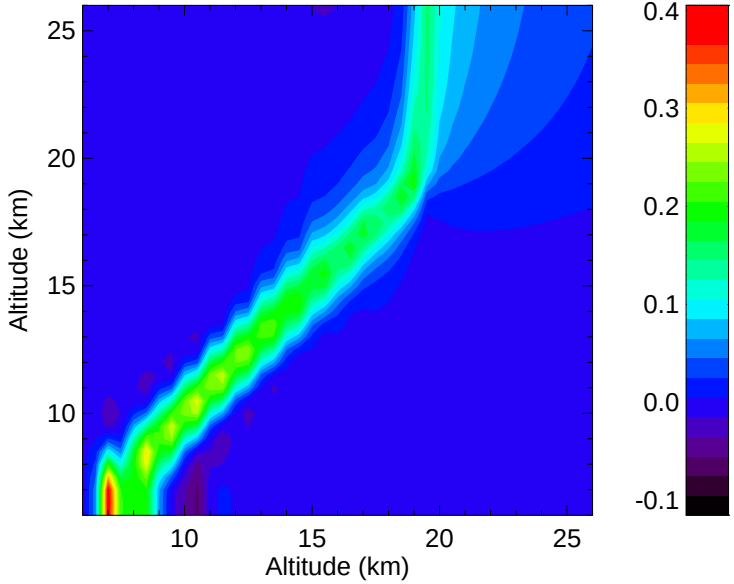

Fig. 7. Averaging kernel for the retrieval of PAN with MIPAS-STR.

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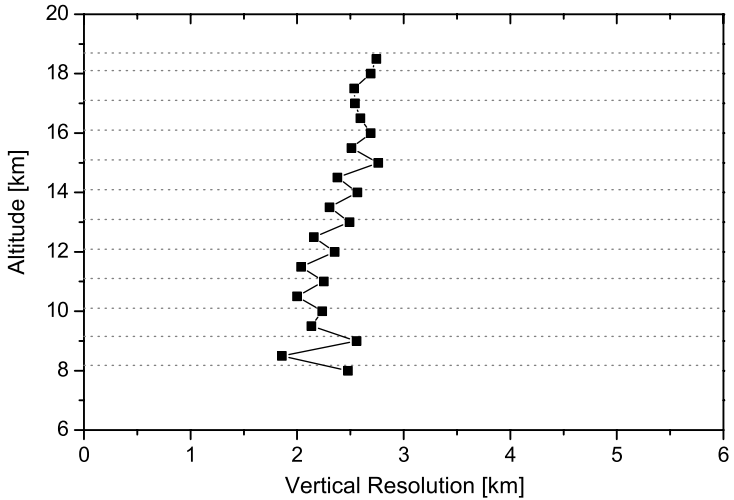

Fig. 8. Achieved vertical resolution of the retrieved PAN profile. The tangent heights are indicated by dotted lines.

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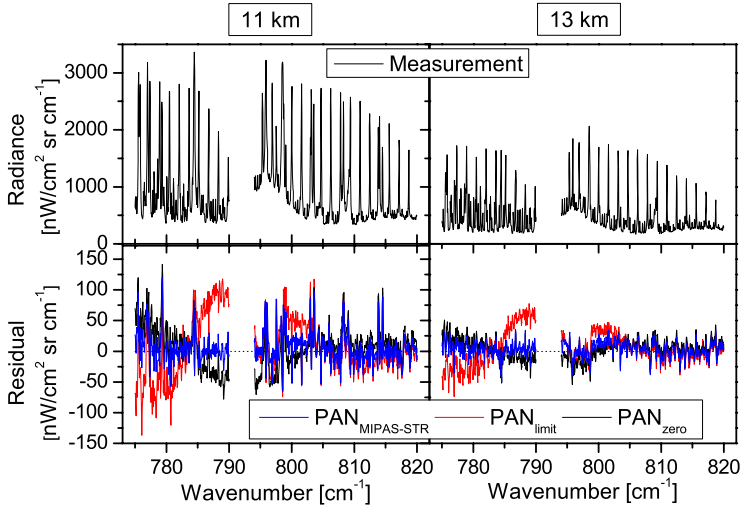

Fig. 9. MIPAS-STR measured spectra in black (top panels) and the residual spectra (forward calculation – measurement) in blue, red and black of RUN_{fit} and the tests RUN_{limit} and RUN_{zero} (lower panels) at tangent heights of 11 km (left panels) and 13 km (right panels).

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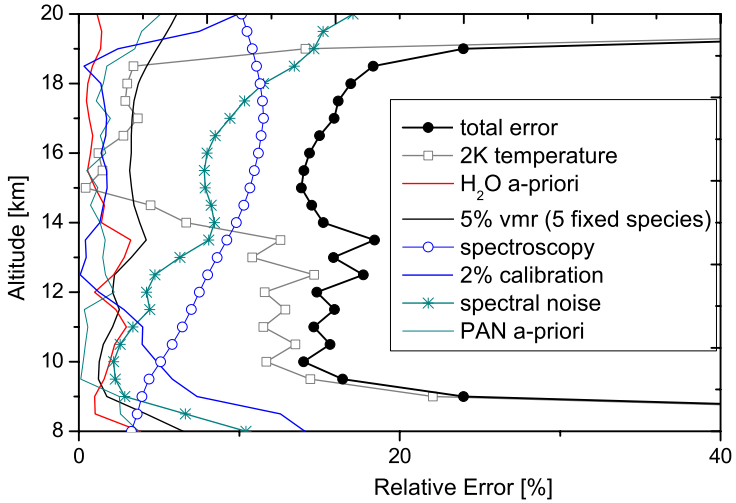

Fig. 10. Estimated total and individual relative errors in $PAN_{MIPAS-STR}$ analysis.

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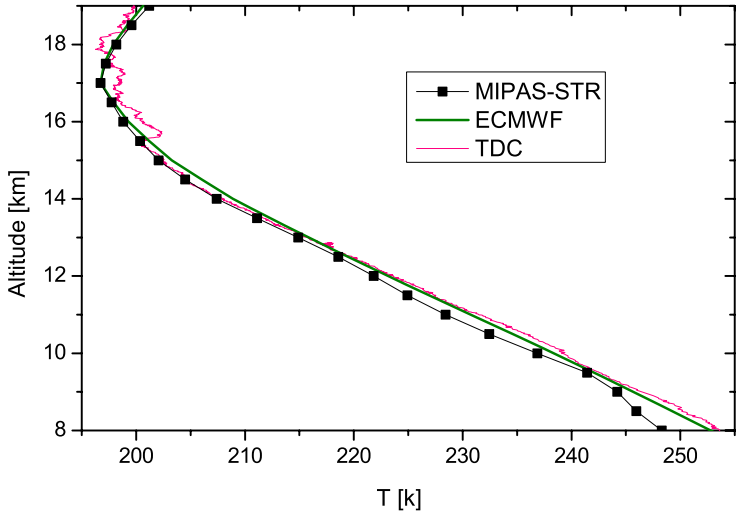

Fig. 11. Comparison of the temperature profile retrieved simultaneously with PAN from MIPAS-STR spectra with that of ECMWF and in situ instrument TDC aboard the Geophysica.

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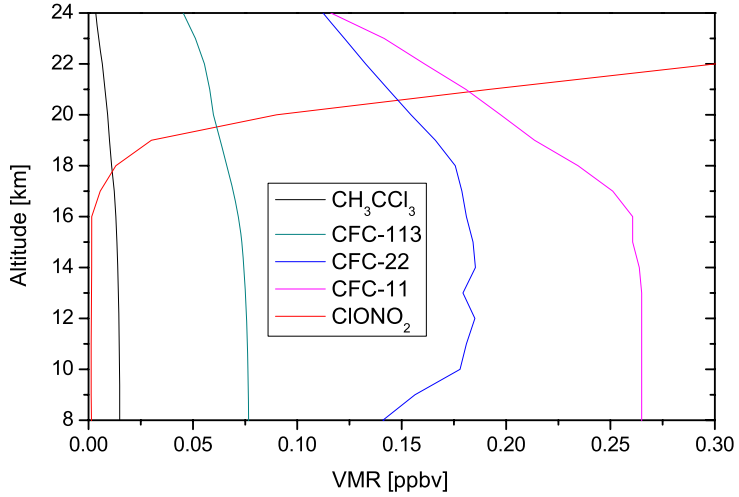

Fig. 12. Vertical profiles of five interfering species determined before the retrieval of PAN.

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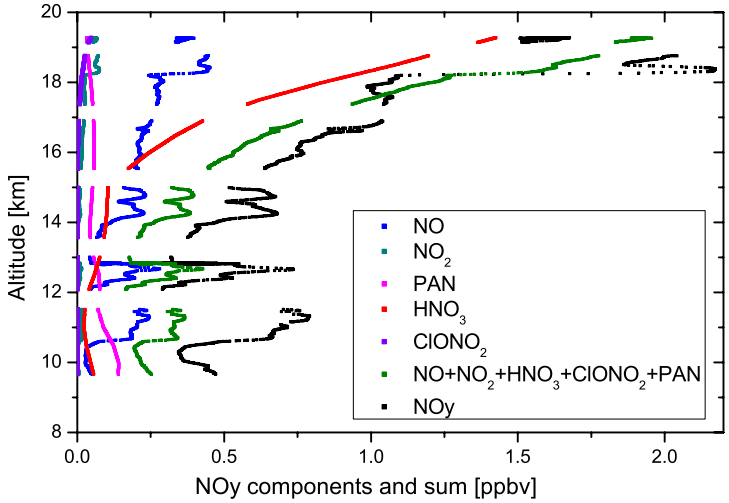

Fig. 13. Comparison of measured NO_y with the profiles of the individual constituents and their sum.

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