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**SCIAMACHY HCHO
and European
isoprene emissions**

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SCIAMACHY formaldehyde observations: constraint for isoprene emissions over Europe?

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Abstract

Formaldehyde (HCHO) is an important intermediate compound in the degradation of volatile organic compounds (VOCs) in the troposphere. Sources of HCHO are largely dominated by its secondary production from VOC oxidation, methane and isoprene being the main precursors in unpolluted areas. As a result of the moderate lifetime of HCHO, its spatial distribution is determined by reactive hydrocarbon emissions. We focus here on Europe, never studied before, and investigate the influence of the different emissions on HCHO tropospheric columns with the CHIMERE chemical transport model in order to interpret the comparisons between SCIAMACHY and simulated HCHO columns. Observed columns present a bias less than 20% on average. The differences are discussed according to the errors on the model and the observations and the remaining discrepancies are attributed to a misrepresentation of biogenic emissions. This study requires the characterisation of: (1) the model errors and performances concerning formaldehyde. The errors on the HCHO columns, mainly related to chemistry and mixed emission types, are evaluated to 2×10^{15} molecule/cm² and the model performances evaluated using surface measurements are satisfactory (~13%); (2) the observation errors that define the needs in spatial and temporal averaging for meaningful comparisons. Perspectives of using SCIAMACHY observations as constraint for biogenic isoprene emissions with an adapted averaging are approached: this new constraint should help to reduce their uncertainties more than 50% in region of intense emissions.

1 Introduction

Volatile organic compounds (VOCs) play an important role in tropospheric chemistry in particular in the production of ozone and organic aerosols, and in odd hydrogen radical and in nitrogen species cycling. They have significant impact for pollution and climate change. Large uncertainties remain in the knowledge of their emissions at continental

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scales as a result of the difficulty in extrapolating local VOC emission measurements to a larger scale. A lot of efforts have been made in order to improve the parameterization of the different effects that influence the emissions, especially for biogenic emissions like the past and present temperatures and radiation, the soil moisture stress, and the age of leaves (e.g. Guenther et al., 2006; Lathièrè et al., 2006; Boissard et al., 2008; Müller et al., 2008, and references therein) but large uncertainties still remain. Satellite observations of formaldehyde (HCHO) have been used as independent constraints for emissions (Palmer et al., 2003, 2006; Abbot et al., 2003; Shim et al., 2005; Fu et al., 2007; Millet et al., 2008).

Formaldehyde is an important intermediate compound in the degradation of the VOCs in the troposphere. Their oxidation by OH leads to the formation of organic peroxy radicals that produce HCHO either directly or through the degradation of higher carbonyls (Atkinson, 1994). The sinks of HCHO are mainly photolysis and reaction with OH. The main source of HCHO in the background troposphere is methane oxidation. However, in the continental boundary layer, oxidation of non-methane VOCs (NMVOCs) dominates over the methane source and can make a large contribution to the HCHO columns. The lifetime of HCHO is variable depending on the OH concentration and on its photolysis. The lifetime of HCHO with respect to OH reaction is ~ 0 at night, varies from 60 h (background global average $[\text{OH}] = 5 \times 10^5 \text{ molecule cm}^{-3}$) to 3 h for $[\text{OH}] = 10^7 \text{ molecule cm}^{-3}$ in photochemically active plume. Overall the lifetime of HCHO is short enough that HCHO column would map the emission fields of its parent VOCs, provided that their reactivity to OH is large enough which is essentially true for unsaturated compounds. The anthropogenic sources of VOCs are the most significant in the urban areas but biogenic sources dominate elsewhere especially during the growing season with the largest contribution coming from isoprene (Palmer et al., 2003, 2006; Abbot et al., 2003; Millet et al., 2008, and figures therein).

A top-down approach for inferring isoprene emissions from HCHO column measurements has been developed by Palmer et al. (2001, 2003) and extensively applied to constrain isoprene emissions in North America using observations from the Global

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Ozone Monitoring Experiment (GOME) UV-visible sounder. In particular the seasonal and interannual variability of the North American isoprene emissions have been studied for the 1995–2001 period (Abbot et al., 2003; Palmer et al., 2006). More recently, Millet et al. (2008) used the Ozone Monitoring Instrument (OMI) to derive isoprene emissions in North America (June–August, 2006) with $1^{\circ} \times 1^{\circ}$ resolution. The use of space-based HCHO data has since been extended to other regions of the world in order to infer biogenic isoprene emissions but also biomass burning and industrial HCHO sources. Fu et al. (2007) used GOME data to constrain NMVOC emissions from east and south Asia, where a complex overlap between high anthropogenic, biogenic and biomass burning emissions occurs. Meyer-Arneke et al. (2005) performed Lagrangian studies over Africa in September 1997 and showed that the major contributions to the measured GOME HCHO columns are biomass burning and biogenic sources. Shim et al. (2005) applied a Bayesian inversion of GOME HCHO column measurements and derived new estimates of biogenic isoprene emissions as well as biomass burning and anthropogenic HCHO sources for eight regions of the world (North America, Europe, east Asia, India, southeast Asia, South America, Africa, and Australia). Except this last study, no detailed studies are available on the possibility of inferring HCHO sources in Europe. This might be explained because the conditions in Europe are less favorable than in other parts of the world. European isoprene emissions are small, 12 Tg C/yr according to Shim et al. (2005), compared to the other regions of the world (106.4 Tg C/yr in Southeast Asia for example). This results in relatively low HCHO columns close to the detection limit of the current space-based instruments able to measure formaldehyde. In addition to weak isoprene emissions, they are well mixed with anthropogenic sources of formaldehyde.

The aim of the present study is to compare for the first time the HCHO tropospheric columns measured and calculated over Europe. The SCAning Imaging Absorption spectrometer for Atmospheric Cartography (SCIAMACHY) sensor aboard the Envisat satellite (Wittrock et al., 2006) and the regional chemical transport model (CTM) CHIMERE (Schmidt et al., 2001) are used. The results are discussed in regards of

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the emission influence accounting for the model and observations errors. This first requires characterizing the performances of the model in reproducing formaldehyde. The chemical scheme included in CHIMERE has been tested with 3 different reference schemes and the simulated surface concentrations of formaldehyde have been compared to surface measurements (Sect. 2). Secondly, the influence of the emissions has to be quantified: a tagging scheme has been included in the CTM and results are presented in Sect. 3. Formaldehyde columns observed with SCIAMACHY are compared to the columns simulated by CHIMERE for two summers (2003 and 2005) with different climatic conditions. The comparisons are discussed taking into account the uncertainties in both the measurements and the simulations and specific averaging are stressed (Sect. 4). Perspectives of using satellite observations of HCHO as constraint of the European isoprene emissions are explored in Sect. 5.

2 CHIMERE model evaluation

2.1 Model description

CHIMERE is an Eulerian multi-scale three-dimensional chemistry-transport model dedicated to air quality issues. It is designed for simulating fields of various pollutants and related species from the urban scale (e.g. Vautard et al., 2001; Menut, 2003; Beekmann and Derognat, 2003; Blond et al., 2003) to the continental scale (e.g. Blond and Vautard, 2004; Besagnet et al., 2004). More recently it has been applied for long-term ozone trends analysis (Vautard et al., 2006), and for diagnostics and inverse modeling of emissions (Deguillaume et al., 2007; Konovalov et al., 2006, 2008; Pison et al., 2007). The model is also used for forecasts of pollutant levels as part of the French national air pollution forecasting system, Prev'Air (www.prevoir.org). The model has been extensively validated with surface and airborne measurements for the different scales (Schmidt et al., 2001; Blond and Vautard, 2004; Menut et al., 2000; Vautard et al., 2003; Blond et al., 2007; Honoré et al., 2008). As an example, the averaged bias for daily

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ozone maxima over Europe is smaller than 10% (Honoré et al., 2008). A detail description of the model is available on the web (<http://euler.lmd.polytechnique.fr/chimere/>).

For this study, the continental version of CHIMERE (Schmidt et al., 2001) vertically extended onto the whole troposphere (Blond et al., 2007) has been used. In this configuration the model covers Western Europe (from 10.5° W to 22.5° E and from 35° N to 57.5° N). The model runs with a horizontal resolution of 0.5° × 0.5° and with 17 vertical layers from the surface up to 200 hPa. The model is driven by the European Center for Medium-Range Weather Forecasts (ECMWF) meteorological analyses. The convective scheme of Tiedtke (1989) is used to rediagnose the convective fluxes. Top and lateral concentrations for 8 species (O₃, NO_x, CO, PAN, CH₄, C₂H₆, HCHO, and HNO₃) are fixed according to the climatological monthly means provided by the MOZART model (Horowitz et al., 2003). Anthropogenic emissions are derived from the EMEP annual totals for 2001 (Vestreng et al., 2004) for NO_x, SO₂, CO, and non methane volatile organic compounds (NMVOC). These emissions are scaled to hourly emissions applying temporal profiles provided by IER (Friedrich, 1997) as described by Schmidt et al. (2001). VOC emissions are distributed into 11 model classes following the mass and reactivity weighting procedure proposed by Middleton et al. (1990): 9 classes are considered for anthropogenic species (n-C₄H₁₀, C₂H₆, o-xylene, C₂H₄, C₃H₆, methylethylketone (MEK), CH₃OH, HCHO and CH₃CHO) and 2 classes for biogenic species (isoprene, terpenes). The biogenic emissions of isoprene, terpenes (represented only by α -pinene) and NO are calculated on-line following the Simpson et al. (1999) and Stohl et al. (1996) methodologies, respectively. The determination of the isoprene and terpene emissions is based on the widely used parameterization of Guenther et al. (1995) that defines the flux as a function of the emission potential, the foliar density for European vegetation (Simpson et al., 1999) and an environmental correction factor that accounts for temperature and radiation dependencies. The emission potential is calculated using specific emission factors for more than 40 tree species and a country-wise tree species distribution. Emission potentials are included also for agricultural land, grass land, and shrubs.

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The complete MELCHIOR chemical mechanism (Lattuati, 1997) is used to perform the simulations of the present study. The photolysis rates are calculated using the troposphere ultraviolet and visible model (TUV) (Mandrovich and Flocke, 1998) and depend on altitude and zenith angle. The mechanism includes a simplified NMHC chemistry and considers a total of 82 gaseous active species and 333 reactions. Deposition of stable intermediates is included. The hydrocarbon species containing three carbons or less are explicitly treated. Larger compounds are represented by lumped species. Biogenic VOCs are represented by isoprene and α -pinene (which represents terpenes). Table 1 summarizes the typical lifetimes and emissions of the NMVOCs emitted in CHIMERE during summertime in Europe. The estimated lifetimes are similar to the lifetimes given by Palmer et al. (2003) in North America but the emissions especially of isoprene are much smaller in Europe than in other regions of the world (Palmer et al., 2003; Shim et al., 2005; Müller et al., 2008).

2.2 Evaluation of the VOC chemistry

The CHIMERE model is used here for the first time to study formaldehyde on a continental scale and has not yet been evaluated for this specific molecule. Thus the reliability of the MELCHIOR chemical scheme with respect to formaldehyde production has to be checked as well as the resulting uncertainties evaluated. We compare calculations of HCHO yields from oxidation of all the VOC emitted in CHIMERE (isoprene, α -pinene, anthropogenic VOCs) using MELCHIOR with the yields derived from well-established chemical mechanisms: (i) the SAPRC99 scheme developed and validated using smog chamber data (for high NO_x conditions) (Carter et al., 2000), (ii) the Master Chemical Mechanism (MCM) (Saunders et al., 2003, Jenkin et al., 2003) which is one of the most detailed schemes available in the literature and often used as a reference, and (iii) the Self-Generated Master Mechanism (SGMM) (Aumont et al., 2005) which is a unique fully explicit chemical scheme, containing up to about 10⁷ reactions (depending on the VOC species treated). Reaction products and rate constants are assigned by an expert system using structure-reactivity relationships. These three

“reference” mechanisms provide a representation of the uncertainties remaining in the NMVOCs oxidation knowledge and allow evaluating how well the chemistry leading to HCHO production is reproduced by MELCHIOR. The yield calculations are made using a time-dependent box model (Aumont et al., 2005; Camredon et al., 2007). We use conditions representative of the ones encountered during summer in Europe: initialization at 09:00 LT for midlatitude summer conditions with 1 ppbv of the VOC studied, 40 ppbv of O₃, 100 ppbv of CO, and either 0.1 or 1 ppbv of NO_x that represent low and high NO_x conditions. The O₃, CO and NO_x concentrations are fixed to their initial values during the simulation time period. Yields are given as formed HCHO molecules per C-atom of a parent VOC. Simulations are performed over 24 h because only the HCHO produced during the first day can be potentially related to the local VOC emissions (Palmer et al., 2006) and then provide pertinent information on the emissions.

Figure 1 shows the time evolution of the HCHO yields obtained in high-NO_x condition (1 ppb). The differences between the three reference schemes (MCM, SAPRC99, and SGMM) are usually within 15% and MCM and SGMM give closer results. However, a large discrepancy between MCM, SGMM and SAPRC is found for the HCHO production from α -pinene oxidation. SAPRC gives a HCHO yield from α -pinene oxidation twice smaller than MCM and SGMM, whereas these 2 schemes are in an agreement of ~20%. Concerning the MELCHIOR mechanism, the HCHO yields derived under high NO_x conditions are globally in fair agreement (<20%) with the reference mechanisms, especially with MCM and SGMM (Fig. 1, Table 2). The yields from C₂H₆, C₃H₆, CH₃CHO, n-C₄H₁₀, and CH₃OH oxidation are in particularly good agreement (<5%) with the three references. The yield from α -pinene oxidation is in a good agreement (~15%) with that derived with MCM and in a very good agreement (<10%) with SGMM. Concerning isoprene oxidation, MELCHIOR leads to a yield slightly smaller: about 10% on average compared to MCM, about 7% compared to SAPRC99 and about 13% compared to SGMM. The performances of MELCHIOR in reproducing the production of HCHO from isoprene oxidation as well as the reactivity of these reactions are very good and better than these of other models: discrepancies larger than 20% are

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noted by Palmer et al. (2006) between GEOS-CHEM and MCM. In low NO_x conditions, the agreement between the chemical mechanisms is degraded compared to the high NO_x conditions (Table 2). However, the agreement remains fairly good for methanol, CH₃CHO, isoprene and MEK. In addition to the disagreement for α -pinene between the reference mechanisms, the largest differences occur for anthropogenic species. However, these species are emitted in regions of high NO_x where the consistency between the different chemical schemes is better (Fig. 1). It is important to note that large uncertainties remain in the oxidation schemes of many VOCs (e.g. α -pinene) and the error estimate gives only an indication of how consistent the different mechanisms are and cannot be an absolute error determination. Moreover, not all the species emitted are explicitly represented in the chemical scheme (family representation). This also introduces uncertainties which are difficult to evaluate.

2.3 Comparison with HCHO surface measurements

In this section formaldehyde simulated by CHIMERE is evaluated using surface measurements gathered within the EMEP ground-based monitoring network (<http://www.nilu.no/projects/ccc/emepdata.html>). Observations are performed on a once- or twice-a-week basis. The sampling time is about 8 h centered at noon. Chemical analysis of the samples is done by reversed phase high-performance liquid chromatography followed by UV detection (Solberg et al., 1995). The precision of the HCHO measurements has been estimated applying parallel sampling and analysis at different stations between 1995 and 2001 (Solberg et al., 1998, 2001a, b, 2005). Three consecutive years of parallel sampling at the Birkenes station (Norway) showed a measurement precision of 5–6% (Solberg et al., 1998). Similar analysis conducted at the French Donon station showed a precision of 7–8% in 1997 (Solberg et al., 1998) and about 13% in 1999 (Solberg et al., 2001a). However, the precision can be much smaller for some cases: an analysis between 1999 and 2001 at the German Waldhof station shows a precision of about 52% (Solberg et al., 2005).

Eight rural monitoring stations (Fig. 2) provided observations on the European do-

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main during summer 2003, the period chosen for comparison. The station type has been characterized using the tagged simulation results (Sect. 3). The different emission contributions are of similar magnitude for the stations of Košetice, Waldhof, Brotjacklriegel, and Schmücke ranging between 15 and 30%. Isoprene oxidation is slightly the largest contribution at Brotjacklriegel and Schmücke whereas terpene oxidation is the largest at Košetice. The station of Zingst, located at a remote site on the Baltic Sea coast, is largely dominated by the background oxidation of methane and the stations at Donon and Peyrusse Vieille, located in remote forested areas, by the isoprene contribution. The HCHO concentrations simulated at the La Tardière station are similarly influenced by methane and isoprene oxidation.

The HCHO surface concentrations observed at each station of the EMEP network between June and August 2003 are compared to the HCHO surface concentrations simulated by the model in Fig. 3. The simulated surface concentrations of HCHO are interpolated at the location of the surface measurements and averaged over the measurement time period. The statistical results are summarized in Table 3. An overall good agreement is obtained between simulations and observations taking the uncertainties of both the simulations and the observations into account: the mean bias over the 3 months is 13%. The daily variability and the mean concentration are rather well reproduced by the model for most of the stations. This indicates that the model captures reasonably well emissions, transport and photochemistry leading to the observed HCHO concentrations. The performance of the model is similar to these of the EMEP model on average (Solberg et al., 2001). There is a tendency for smaller biases and larger correlation coefficients for the three French sites with small anthropogenic contributions. On the contrary simulations underestimate surface formaldehyde for the station of Zingst (40%). Significant discrepancies for a coastal site are not surprising. Honoré et al. (2008) stressed a general problem for the score of the model at coastal sites for simulating daily ozone maxima with the continental version of CHIMERE. A large bias appears for 2 stations (Waldhof and Brotjacklriegel) with simulations larger by a factor between 2 and 3. As contributions from different sources are well equili-

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brated at these two sites, it is not obvious which process is responsible for that difference. Some well admitted model errors can be cited as a potential explanation for that difference: uncertainties in different types of VOC emissions (see below), uncertainties in the chemical scheme, uncertainties in OH concentrations affecting the HCHO formation and loss, uncertainties of the meteorology affecting transport patterns.

3 Contribution of the emissions to the HCHO tropospheric columns

The aim of the present section is to evaluate how the emissions and the tropospheric columns of HCHO are connected over Europe. The mean spatial distributions of the VOC emissions considered in CHIMERE in summer 2003 are represented in Fig. 4. The anthropogenic emissions dominate the other emissions mainly in the North-West of Europe (North of France including Paris area, Belgium, Netherlands, West of Germany and England), in the Po Valley and along the East coast of Italy. Anthropogenic and biogenic emissions are of comparable magnitude over Spain. Isoprene emissions dominate in the other European regions except in the Czech Republic and Slovakia where the terpene emissions dominate.

In order to identify the regions where the impact of the different emissions and their variations would be detectable, we included different chemical tagging schemes into the CHIMERE model using a similar technique as Pfister et al. (2008). This enables the contribution of different formaldehyde production pathways to the tropospheric column to be tracked separately. Four pathways are considered: (i) methane oxidation (background contribution), (ii) isoprene oxidation, (iii) terpene oxidation and (iv) oxidation of the anthropogenically emitted VOC. It is verified that the tagging procedure does not alter results for the original species. Only directly emitted anthropogenic VOCs and not those produced in addition in the different stages of the oxidation chains leading to HCHO are considered in the experiment. Formaldehyde produced by emitted CH_3CHO , CH_3OH and MEK is then not tagged here. This simplification is reasonable because the primary emissions of these 3 species are negligible compared to the other

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anthropogenic VOCs (Table 1). However, methanol has a large biogenic source that is not accounted for in CHIMERE: the biogenic part is more than one order of magnitude larger (7.4×10^{10} molecule/cm²/s, Granier et al., 2005) than the anthropogenic part (0.3×10^{10} molecule/cm²/s). According to the HCHO yield from the methanol oxidation, the potential contribution to the non-background part of the HCHO column is about 10%. The relatively long lifetime of methanol (Table 1) suggests that the formaldehyde produced by the biogenic methanol emitted in the model domain would be widely dispersed over the domain and would not provide any information on the emissions. Thus neglecting biogenic methanol introduces an error on the simulated tropospheric columns of HCHO of the order of 10%.

The 2003 summer means of the different contributions are presented in Fig. 5. The results are presented at 10:00 LT, i.e. at the overpass time of SCIAMACHY (see Sect. 5). The background contribution represents about 55% of the HCHO tropospheric columns on average but can be as low as a few percent in some grid cells and reaches up to 85% in some others, especially over sea. The mean isoprene contribution to the HCHO tropospheric column is about 17% and 21% with and without over-sea columns, respectively. Isoprene oxidation is the largest source of formaldehyde over the southwestern half of France, over Greece and North Africa. It also dominates contributions from other emissions over Poland and the Balkans. Note that the impact of the isoprene emissions on the HCHO columns is mainly localized above the source regions. The spatial correlation (r^2) between the emissions and the isoprene-tagged formaldehyde columns is 64%. This reflects that the smearing effect is relatively limited. The terpene emissions can contribute to up to 25% of the HCHO column but their mean contribution is rather low (8% on average). The maxima of the contribution are also localized close to the main source regions but the spatial correlation of the tagged columns with the emissions is smaller (45%) and most likely results from the lower contribution to the HCHO as compared to that of isoprene. The mean contribution of the anthropogenic emissions to the HCHO column (11% on average) is also smaller compared to the isoprene contribution. The spatial correlation between the anthropogenic emissions and

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the tagged columns is also reduced (44%) compared to isoprene and only the unsaturated species significantly contribute to this correlation. Alkane emissions represented by n-butane are too unreactive to be oxidized in the same grid cell (Table 2). A large contribution (up to 40%) is observed above Northern Italy with the largest columns localized above the most intense source regions. This implies that the influence of the most reactive species in this case like C_2H_4 , C_3H_6 or o-xylene (Fig. 1) on HCHO is being observed. These molecules are also emitted with the same order of magnitude in the North-West of Europe but with a much smaller impact on the HCHO columns. This behavior is attributed to two factors: (i) the longer residence time of pollutants over the Po Valley (lower winds), (ii) the larger oxidative capacity of the air in this region, which results from larger amounts of actinic radiation, relatively high ozone and water vapour concentrations (increased OH production) and lower radical losses due to lower NO_x levels. Finally, the sum of all the contributions (background + emissions) is shown in Fig. 5. This map shows the influence of the boundary conditions on the HCHO tropospheric column (from the difference to 100%). In agreement with the dominant winds in Europe, the boundary conditions have a significant impact mainly close to the western and the northern boundary of the domain. The contribution ranges from 20 to 30% in the boundary regions to more than 70% in the extreme West of the domain. Ireland and the North of United Kingdom are the most affected areas. The contribution decreases rapidly towards the interior of the domain according to the short formaldehyde lifetime of several hours.

It is also worth noting that the temporal and spatial variability of the HCHO tropospheric columns is driven on average over the European domain by isoprene variability, followed by the background variability. The influence of anthropogenic VOCs and terpenes is limited except in regions with large emissions. The temporal variability (during summertime) is estimated by the standard deviation for the mean tagged column over the European domain and then likely underestimated. The standard deviation value is 0.37×10^{15} molecule. cm^{-2} for isoprene-tagged columns, 0.27×10^{15} molecule. cm^{-2} for background-tagged columns, and 0.10 and

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0.17×10¹⁵ molecule.cm⁻² for anthropogenic- and terpene-tagged columns, respectively. The spatial variability is estimated by the standard deviation of the European tagged columns averaged over the June–August period. The standard deviation values are larger than for the temporal variability with a value of 1.11×10¹⁵ molecule.cm⁻² for isoprene (1.01, 0.51 and 0.42×10¹⁵ molecule.cm⁻² for background, anthropogenic and terpenes, respectively). As a conclusion, methane oxidation is the dominant source of formaldehyde (considering the total tropospheric columns). The amount of formaldehyde produced by isoprene oxidation is significantly larger than formaldehyde produced by the other emitted VOC and dominates the HCHO NMVOC production except for very localized areas (e.g. North of Italy). Isoprene is also the main driver of the temporal and the spatial variability.

4 SCIAMACHY/CHIMERE comparison over Europe in summer 2003

4.1 Data description

SCIAMACHY is an instrument operating onboard the sun-synchronous Envisat satellite since 2002 which has an equator crossing time of 10:00 a.m. LT. This instrument is a spectrometer designed to measure sunlight transmitted, reflected and scattered by the earth's atmosphere or surface in the ultraviolet, visible and near infrared (Bovensmann et al., 1999). The maximum scan width in the nadir-view is 960 km and global coverage is achieved within 6 days. The horizontal resolution in the nadir mode is 30×60 km².

The retrieval of vertical columns of HCHO is based on the DOAS technique that relies on the separation of narrow band absorption signatures from broad band absorption and scattering features. The retrieval consists in the determination of the slant column density of the considered species and its conversion to a vertical column amount by applying an air mass factor (AMF). This accounts for the path of light through the atmosphere and takes the vertical profiles of scattering and absorbing species into account. The spectral region of 334–348 nm was selected for the retrieval to avoid any corre-

lation with an instrument grating polarization structure around 360 nm. In comparison to HCHO measurements from GOME (Global Ozone Monitoring Experiment, e.g. Wittrock, et al., 2000) this leads to a reduced signal-to-noise-ratio. The AMF calculation uses three standard profiles based on observations reported in literature, which are assigned to the individual pixels and times based on external information (Wittrock, 2006). Only ground scenes having less than 20 percent cloud cover are considered. Prior to conversion to vertical columns, the slant columns were normalized by assuming a mean value of 3.5×10^{15} molecule.cm⁻² in the region between 180° W and 160° W in agreement with climatological values from the MOZART (Horowitz et al., 2003) and the LMDz-INCA (Hauglustaine et al., 2004) models. This normalization is necessary to compensate for offsets introduced by the solar reference measurements (Richter and Burrows, 2002) and interferences by other absorbers. The uncertainty of HCHO vertical column retrieved from SCIAMACHY measurements depends mainly on the fitting accuracy. For a single pixel (30×60 km² footprint) it is about 10¹⁶ molecule/cm². By averaging both in space and time this can be significantly reduced (see section below). The surface spectral reflectance, the assumption for the aerosol vertical distribution and its optical thickness, and the assumed vertical distribution of formaldehyde in the lowermost troposphere contribute to the uncertainty of the AMF calculation in the retrieval. This error is strongly scene-dependent but varies typically in the range of 10 to 30% of the vertical column. Systematic error sources are the temperature dependence of the HCHO cross section and inaccuracies of its absolute value, spectral interference to other trace gases and the normalisation of the slant columns applying a reference sector. These sources sum up to about 3×10^{15} molecule/cm². A detailed discussion of the data evaluation can be found in Wittrock, 2006. SCIAMACHY data are regridded on a 0.5°×0.5° grid for an initial use to match the model resolution.

4.2 Results

The comparison of daily SCIAMACHY observations and simulations above Europe is not reasonable because the mean European column of formaldehyde

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($\sim 7 \times 10^{15}$ molecule/cm²) is smaller than the error (random error for one pixel about 10^{16} molecule/cm², see above). Thus, it is necessary to average the observations either temporally or spatially (or both) to reduce the errors. In a first step, we choose to temporally average the observations on a seasonal (summer months) and a monthly basis. The random error is reduced by the square-root of the number of SCIAMACHY overpasses during the time period considered. On average the number of overpasses is about 6.25 per month and 18.75 per season for one pixel. A detailed calculation for each pixel indicates that the total error decreases down to 3.8×10^{15} molecule/cm² on average when observations are averaged over the 3 summer months and down to 5.4×10^{15} molecule/cm² when the average is made over one month. This latter error is still large compared to the mean HCHO column. In order to have comparable error level over 3 months and over one month, the observations have been regridded on a $1^\circ \times 1^\circ$ grid for the monthly means.

The summer and monthly tropospheric formaldehyde columns measured by SCIAMACHY in 2003 and 2005 are displayed in Fig. 6 (middle and lower panels) on a $0.5^\circ \times 0.5^\circ$ grid for the JJA period and $1^\circ \times 1^\circ$ for individual months. For 2005 that is a rather normal year concerning climatic conditions the HCHO columns observed by SCIAMACHY decrease from June to August (Fig. 6 and Table 4) but they are too close to the detection limit of the instrument to be unambiguously compared to model simulations, especially in July and August. In 2003, the mean European columns (7.33×10^{15} molecule/cm²) for the JJA period is 50% larger than in 2005 (Table 4). The exceptionally hot temperatures that had persisted during all the summer with a maximum at the beginning of August implied much larger biogenic emissions (Curci et al., 2008). Combined with stagnating anticyclonic conditions, this leads to much larger amount of HCHO produced. The temporal evolution of HCHO columns is different from 2005: even if the columns observed in July are smaller than those observed in June, unusually large columns are measured in August during the heat wave period (Fig. 6 and Table 4). The continental columns are generally larger than the oceanic columns in agreement with the significant part of continental emissions leading to HCHO produc-

tion. The oceanic columns are quite similar between 2003 and 2005 (Table 4) except for August when heat wave influence was the largest. Moreover, the spatial variability (standard deviation) of the columns is also larger in continental areas than over ocean (Table 4). One expects that the increased variability is mainly due to an increased variability of the sources over land. However, some observations have to be taken with caution especially the monthly average also for 2003 because observed column values are in the range of their total errors for example for Spain in June and July. Some of the large columns observed over sea are also questionable because they might be artifact due to interfering contributions that can perturb the HCHO retrieval (de Smedt et al., 2008). However, large wild fires occurred in August 2003 in Portugal and Spain and consequently their emissions – not included in CHIMERE – could explain the large columns observed in the Gascony gulf (e.g. Fig. 12 in Tressol et al., 2008).

The formaldehyde columns observed by SCIAMACHY in summer 2003 are compared to the columns simulated by CHIMERE (upper panel) in Fig. 6 and the statistical results of the comparison are summarized in Table 4. The CHIMERE fields are taken at the time and location of the SCIAMACHY measurements for the comparison. Most of the main continental spatial structures observed by SCIAMACHY are qualitatively reproduced by the model (Fig. 6 – JJA period). Unusually large formaldehyde values during August are also simulated by the model in agreement with strong heat wave conditions encountered in Europe for this period. However, the magnitude of these increased columns is less in the model and the spatial distribution cell by cell is not well reproduced. A small tendency of the model to underestimate formaldehyde is also found for June but not for July (Table 4). The mean biases for individual months are smaller than 25% and are clearly within the measurement and model error limits. The relative RMSE (root mean square of the error) of 40–50% is large but is of the order of the SCIAMACHY uncertainties. The differences between the observations and the simulations are more important for the oceanic columns (Table 4) but these columns have to be considered with caution as mentioned previously. If one focuses on the continental part, the differences are significantly reduced: the mean bias for

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the JJA period decreases by a factor of 2, RMSE and spatial correlations are rather similar. The main differences between the observations and the simulations occur in the South-East of France, in North Africa and more generally for the latitude larger than 50° N whatever the time period considered. The discrepancy in North Africa is not surprising because the biogenic emissions used are more uncertain because tree species distribution is extrapolated from a typical Mediterranean one taken from the Greek one. The discrepancies at the northern latitudes are difficult to explain. They could be partly related to emissions but also to a degradation of the SCIAMACHY measurement quality when going north that is not yet clearly understood. Note that the model to observation differences noted over North-Eastern Germany have different signs when tropospheric columns or surface values are compared (positive bias for simulated surface values at Waldhof, negative bias for tropospheric columns). In addition, as discussed in Sect. 2.3, the simulated columns in the North of Europe (UK, Poland) are the most affected by the boundary conditions. The use of climatological monthly means to force the model at the boundaries may also explain some of the differences in these regions. However, the small spatial correlations obtained when model and observations are compared cell by cell (Table 4) suggest that such comparisons are not the most satisfying because significant noise remains in the measurements. Additional spatial averaging on consistent formaldehyde column fields helps to reduce the measurement noise and to have more pertinent comparison in regards of the emissions. Rather to degrade again the horizontal resolution without any relations with the source regions, we chose to define subdomains for their consistency in terms of biogenic emissions influence onto HCHO and for the specific differences noted between SCIAMACHY and CHIMERE (e.g. South-East of France or Poland). Nineteen subdomains have been finally selected (Fig. 7). The SCIAMACHY and CHIMERE formaldehyde columns have been spatially and temporally averaged over each subdomain and compared. The total error on the SCIAMACHY columns is now reduced to the systematic error (3×10^{15} molecule/cm²): the random error contributes with less than 1%. The total error corresponds to 30 to 70% of the tropospheric columns depending on

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the subdomain.

The comparison results between SCIAMACHY and CHIMERE mean columns are shown in Fig. 8 for the JJA period. The agreement between SCIAMACHY and CHIMERE is better than 20% for the majority of the subdomains. The differences are the largest in the most northern domains (5, 6, 17, and 19) with SCIAMACHY larger than the model. On the contrary, the observed columns are smaller (>20%) than the simulated ones in subdomains 2, 11 and 12. Subdomains 11 and 12 are also slightly impacted by the boundary conditions. Uncertainty in biogenic emissions is larger for domain 12 (North Africa) as seen previously. Thus on one hand, the differences between observations and simulations are within the SCIAMACHY errors for almost all the subdomains, on the other hand, they can be explained by the large uncertainties especially in biogenic emissions (see below for more detailed discussion).

5 Potential gain of satellite data use for emission improvement

In this section, the potential of satellite data (here SCIAMACHY) as potential constraint for inferring emission information is discussed with a specific focus on the limitations implied by the instrument and model performances and the mixing of different emission types. The detection limits in terms of isoprene emissions discussed in the following are estimated as in Millet et al. (2008): the limiting HCHO column is divided by a typical HCHO column lifetime (~2.5 h, Table 1) multiplied by the ultimate HCHO molecular yield from isoprene oxidation (2.5). The intrinsic limitation in isoprene emission detection from HCHO column observation is linked with the error in the variability of the other emissions and especially the temporal variability if the objective is to derive gridded emissions. The temporal variability of the other emissions is approximated by the variances of the mean tagged HCHO columns resulting from methane, terpenes (α -pinene), and anthropogenic species for the European domain. Considering an error of 10% for methane concentrations used in CHIMERE, of 40% for anthropogenic emissions (i.e. Beekmann and Derognat, 2003) and of a factor of 3 for terpene emissions

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(Simpson et al., 1999), the square sum of the errors on the variances leads to a limiting HCHO column of 0.25×10^{15} molecule/cm² that corresponds to an isoprene emission limit of 1.13×10^{10} molecule/cm²/s. However, this detection limit in terms of HCHO columns (0.25×10^{15} molecule/cm²) is much smaller than the SCIAMACHY uncertainties. These latter are then the main limiting factors of isoprene emission detection in Europe. However, if one considers the admitted uncertainties in the biogenic emissions that range between a factor of 2 and 3 (Simpson et al., 1999), and the errors in the observations and the simulations, one can estimate the improvement of the emissions knowledge brought by the satellite observation using inverse modeling theory (i.e. Rodgers, 2001). In this formalism, the a posteriori error on the inversed emissions σ is linked to the a priori error on the emissions σ_a , to the observation errors σ_{obs} , to the model errors σ_{mod} , and to the sensitivity of the HCHO columns to isoprene emissions k :

$$\frac{1}{\sigma^2} = \frac{1}{\sigma_a^2} + \frac{k^2}{\sigma_{\text{obs}}^2 + \sigma_{\text{mod}}^2} \quad (1)$$

Considering the inventory used in CHIMERE, a factor of 2 is considered for σ_a . The observation errors (σ_{obs}) are equal to the systematic errors (3×10^{15} molecule/cm²) if subdomains are considered. The model errors have been discussed in Sect. 2.3 concerning the chemistry: an error of 10% on the HCHO production from isoprene oxidation is reasonable. In addition it is necessary to account for the terpene contribution that is largely uncertain as well as the emissions. Considering the mean contribution to the column (8%, Sect. 3) and an uncertainty of a factor of 3 for the emissions, the errors linked to terpenes can be estimated to 1.75×10^{15} molecule/cm². Combined with the uncertainty on the HCHO production from isoprene oxidation and the uncertainty resulting of the non-consideration of biogenic methanol, the model error used (σ_{mod}) is 2×10^{15} molecule/cm² on average. Note that the contribution of the uncertainty on the anthropogenic part of the column is negligible compared to the other error sources and thus not accounted for in the calculation. The sensitivity of the HCHO columns

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to isoprene emissions k is estimated using the isoprene-tagged columns and defined against a multiplying factor of the emissions (without units). It ranges from 4.65×10^{15} (subdomain 3) to 0.70×10^{15} molecule/cm² (subdomain 19) and leads to a reduction of about 61% and 6% of the a priori emission uncertainties if current SCIAMACHY data are used. Thus, it is expected that for sensitive grid-cells, satellite HCHO column observations can significantly reduce uncertainty in European isoprene emissions.

6 Conclusions

For the first time, comparisons of HCHO tropospheric column measurements for June–August 2003 from SCIAMACHY aboard the Envisat satellite with simulated columns with a specific focus on Europe are presented. The columns observed in 2003 were well above the detection limit of the instrument and we showed that the most important observed spatial structures in HCHO columns are reproduced by the model (mean bias <20%) but that differences likely related to biogenic isoprene emissions uncertainties remains. The discussion of the results required:

(1) the evaluation of the model performances and the errors (mainly related to HCHO chemistry). Comparisons with surface measurements show that the model succeeds to reproduce the basic photochemistry leading to the observed HCHO concentrations for most of the sites. Unfortunately the small number of sites operating in-situ formaldehyde measurements does not allow a complete model evaluation. On the other hand, no strong inconsistencies concerning HCHO chemistry in MELCHIOR mechanism have been revealed. The error on the HCHO production from isoprene oxidation is estimated to 10%. Addition of other error contributions (terpenes, methanol) leads to a model error of $\sim 2 \times 10^{15}$ molecule/cm².

(2) the definition of regions of interest in Western Europe in regards of the isoprene emissions. This was achieved by implementing tagged chemical schemes model in order to infer the contribution of the different sources of formaldehyde in the European tropospheric column and then the most sensitive to isoprene.

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(3) The observations errors that also controlled the definition of the regions of interest stressing the needs of specific spatial and temporal averaging over coherent HCHO sources.

Finally, perspectives of using satellite observations as constraint of isoprene emissions in Europe are approached. We evaluated for example that SCIAMACHY observations used as constraint could reduce the errors on the emissions more than 50% in well chosen regions.

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Table 1. NMVOCs lifetime and emissions over Europe in summer.

Species	Lifetime ^a	Emissions (10 ¹⁰ molecule/cm ² /s)	Species	Lifetime ^a	Emissions (10 ¹⁰ molecule/cm ² /s)
Isoprene	32 min	6.05	C ₃ H ₆	2 h	0.40
α -pinene	47 min	3.05	MEK	1 day	0.28
n-C ₄ H ₁₀	22 h	3.29	CH ₃ OH	2.5 days	0.27
C ₂ H ₆	9 days	0.51	HCHO	2.5 h	0.19
o-xylene	4 h	0.47	CH ₃ CHO	3 h	0.06
C ₂ H ₄	6 h	0.40			

^a VOC lifetime under summer midmorning conditions are calculated assuming $[\text{OH}]=5\times 10^6$ molecule cm⁻³, $[\text{O}_3]=10^{12}$ molecule cm⁻³ (~40 ppbv), a temperature of 298 K and a solar zenith angle of 32.4°.

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Table 2. One-day yields of formaldehyde per C reacted from the oxidation of the parent VOCs emitted in CHIMERE for 1 and 0.1 ppbv of NO_x.

Species	1 ppbv NO _x				0.1 ppbv NO _x			
	MCM	SAPRC	SGMM	MELC	MCM	SAPRC	SGMM	MELC
Isoprene	0.47	0.45	0.48	0.42	0.30	0.27	0.25	0.28
α -pinene	0.21	0.11	0.16	0.17	0.07	0.04	0.03	0.08
n-C ₄ H ₁₀	0.15	0.16	0.14	0.15	0.02	0.05	0.02	0.03
C ₂ H ₆	0.05	0.05	0.05	0.05	0.01	0.03	0.01	0.01
O-xylene ^a	0.23	0.18	–	0.26	0.09	0.06	–	0.14
C ₂ H ₄	0.83	0.91	0.88	0.98	0.50	0.79	0.53	0.88
C ₃ H ₆	0.61	0.66	0.61	0.64	0.36	0.50	0.29	0.49
MEK	0.22	0.16	0.21	0.20	0.07	0.06	0.06	0.06
CH ₃ OH	0.38	0.39	0.39	0.41	0.27	0.27	0.26	0.27
CH ₃ CHO	0.49	0.48	0.49	0.48	0.30	0.28	0.30	0.30

^a Oxidation schemes of aromatics (here o-xylene) are not available from SGMM.

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Table 3. Statistics inferred from the comparison between surface observations and simulations for the June-July-August 2003 period at each EMEP stations: obs and sim represent the mean surface concentrations of HCHO observed and simulated respectively; R is the correlation coefficient; n the number of comparison points, the bias is given as sim-obs; and RMSE is the root mean square of the error between the observations and the simulations.

Station	obs	sim	R	n	Bias		RMSE	
	($\mu\text{g}/\text{m}^3$)	($\mu\text{g}/\text{m}^3$)			($\mu\text{g}/\text{m}^3$)	(%)	($\mu\text{g}/\text{m}^3$)	(%)
CZ03	2.40	2.41	0.47	26	0.007	0.3	1.03	43
DE02	1.18	2.15	0.79	25	0.98	83	1.17	99
DE05	1.01	2.90	0.53	26	1.89	187	2.04	202
DE08	2.84	2.64	0.36	25	-0.20	-7	1.37	48
DE09	1.99	1.19	0.79	26	-0.80	-40	0.91	46
FR08	3.76	4.12	0.80	13	0.37	10	1.20	32
FR13	2.93	2.78	0.61	13	-0.16	-5	1.87	64
FR15	2.27	2.10	0.86	13	-0.17	-7	0.64	28
all	2.14	2.43	0.50	167	0.29	13	1.84	86

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Table 4. Statistics derived from the comparison of HCHO tropospheric columns (10^{15} molecule/cm²) observed by SCIAMACHY and simulated by CHIMERE for summer 2003 and 2005. The statistics are performed on the total domain, and on the land and sea columns, respectively.

year	n	2003					2005				
		obs ^c	sim	bias ^d	RMSE ^d	R ^e	obs ^c	sim	bias ^d	RMSE ^d	R ^e
All											
JJA ^a	2990	7.33 (2.77)	6.16	−1.17	2.99	0.39	4.91 (2.51)	5.70	0.79	2.93	0.11
Jun ^b	736	7.30 (3.25)	5.77	−1.53	3.50	0.38	6.61 (3.16)	5.64	−0.97	3.43	0.20
Jul ^b	736	5.99 (3.38)	6.23	0.24	3.29	0.33	4.92 (3.13)	6.00	1.08	3.58	0.11
Aug ^b	736	9.18 (3.96)	6.87	−2.31	4.50	0.37	3.32 (3.18)	5.49	2.18	3.98	0.17
Land											
JJA ^a	1936	7.34 (2.91)	6.73	−0.61	2.93	0.38	4.47 (2.55)	6.10	1.63	3.27	0.12
Jun ^b	526	7.41 (3.48)	6.21	−1.20	3.48	0.41	6.44 (3.25)	5.96	−0.48	3.35	0.24
Jul ^b	526	6.03 (3.50)	6.65	0.62	3.52	0.28	4.56 (3.14)	6.39	1.83	3.86	0.13
Aug ^b	526	9.16 (3.97)	7.37	−1.79	4.30	0.37	2.98 (3.08)	5.75	2.77	4.28	0.15
Sea											
JJA ^a	1054	7.31 (2.50)	5.13	−2.18	3.10	0.50	5.72 (2.21)	4.95	−0.76	2.17	0.44
Jun ^b	210	7.02 (2.58)	4.69	−2.33	3.54	0.26	7.03 (2.89)	4.83	−2.19	3.61	0.23
Jul ^b	210	5.89 (3.09)	5.18	−0.59	2.61	0.59	5.82 (2.89)	5.04	−0.78	2.76	0.41
Aug ^b	210	9.23 (3.94)	5.63	−3.60	4.96	0.49	4.16 (3.27)	4.85	0.69	3.08	0.40

^a resolution $0.5^\circ \times 0.5^\circ$

^b resolution $1^\circ \times 1^\circ$

^c The standard deviation of the mean European HCHO columns observed by SCIAMACHY is indicated between parentheses.

^d The bias and the RMSE are given in absolute value.

^e Spatial correlation

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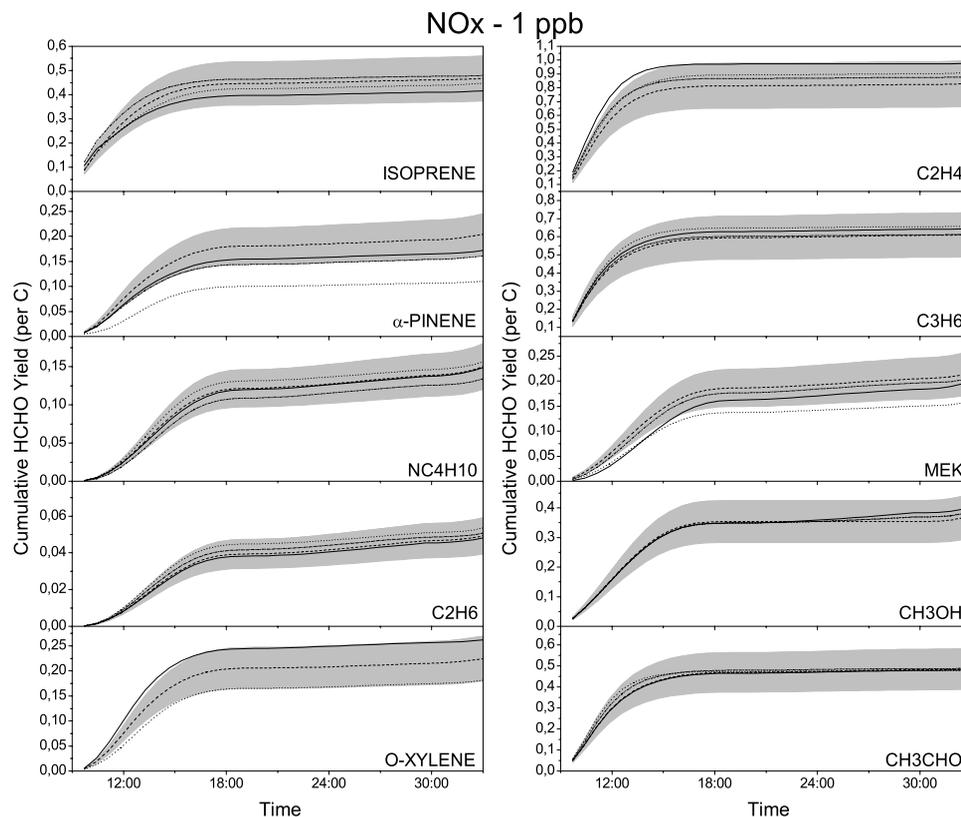


Fig. 1. Cumulative HCHO yields per carbon from the oxidation of the VOCs emitted in CHIMERE compared for different chemical mechanisms under high NO_x conditions (1 ppb). Solid line: MELCHIOR scheme; dashed line: MCM scheme; dotted line: SAPRC scheme; dashed-dotted line: SGMM scheme. The shaded area corresponds to $\pm 20\%$ deviation around the yield simulated with MCM.

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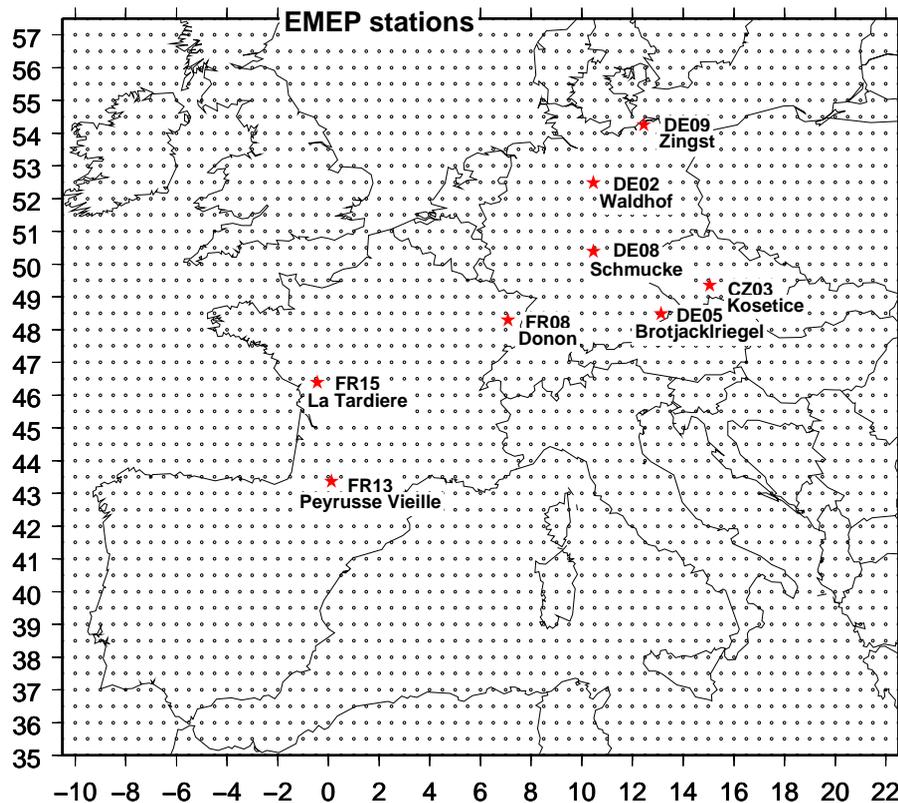


Fig. 2. Location of the EMEP stations used for the comparison of surface HCHO concentration.

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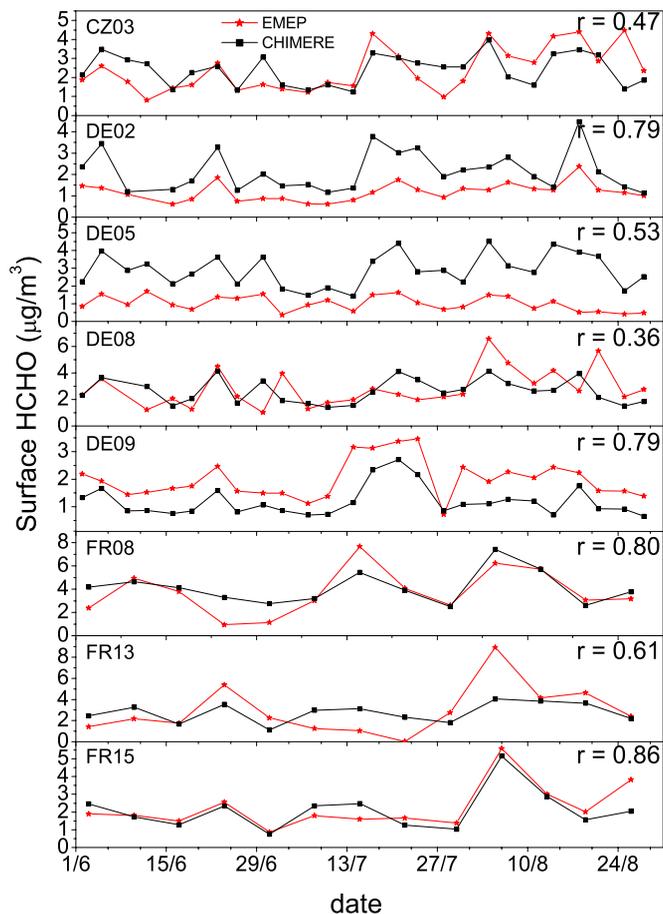


Fig. 3. Timeseries of surface HCHO concentrations during summer 2003 as measured at the 8 EMEP stations and simulated by CHIMERE.

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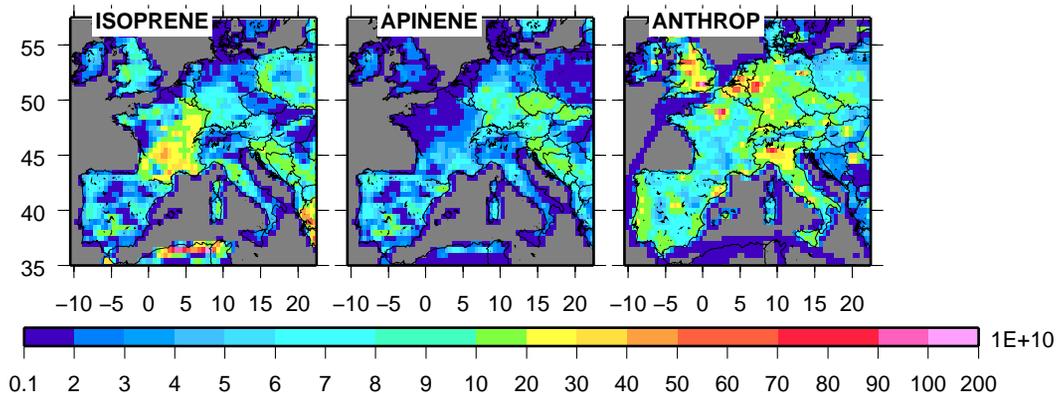


Fig. 4. Statistics derived from the comparison of HCHO tropospheric columns (10^{15} molecule/cm²) observed by SCIAMACHY and simulated by CHIMERE for summers 2003 and 2005. The statistics are performed on the total domain, and on the land and sea columns, respectively.

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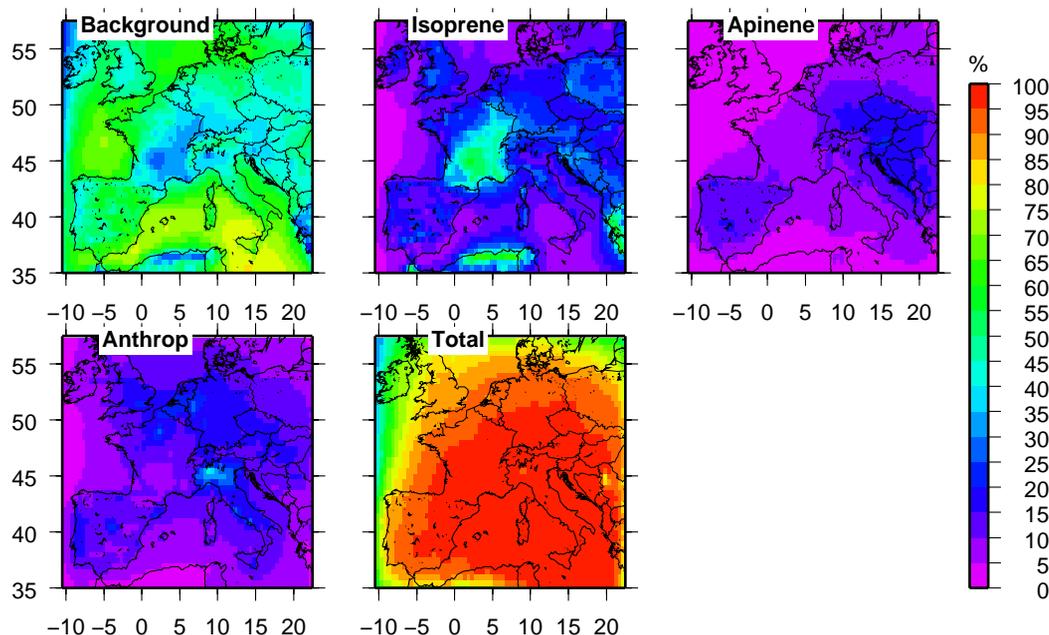


Fig. 5. Contribution (%) to the tropospheric column of HCHO from methane (background), isoprene, α -pinene, and anthropogenic VOC oxidation. The sum of the different contribution is also presented in order to provide information on the boundary condition influence (difference to 100%).

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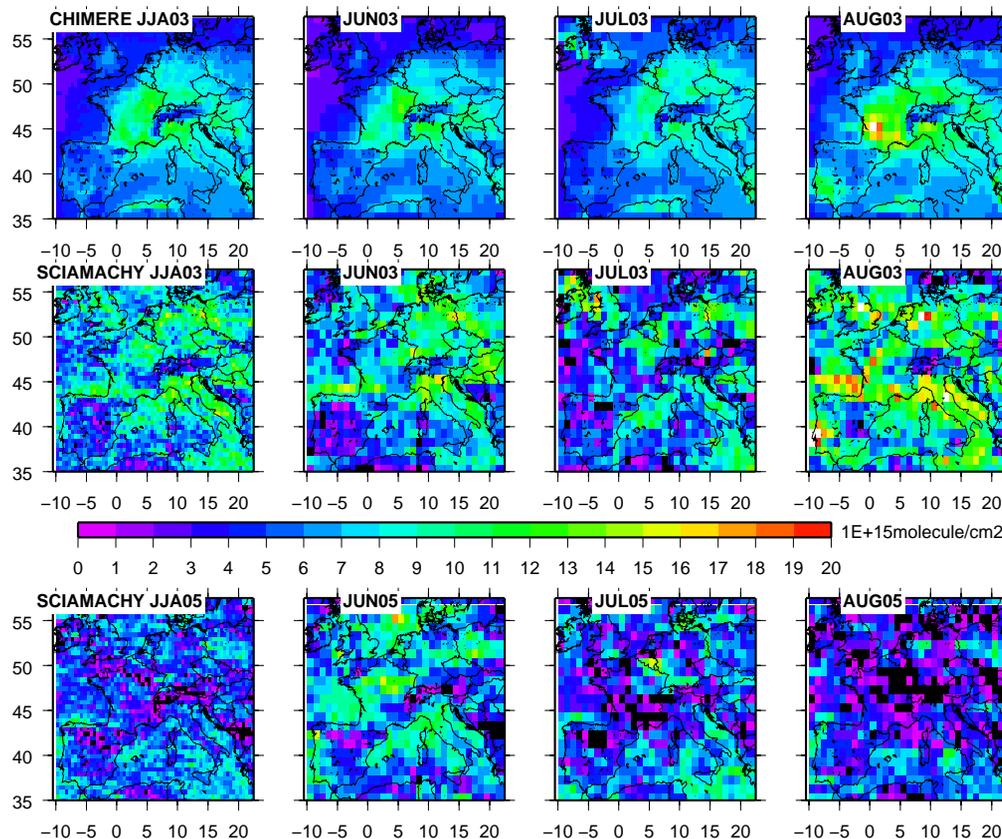


Fig. 6. HCHO tropospheric columns simulated by CHIMERE in summer 2003 (top panel), observed by SCIAMACHY in summer 2003 (middle panel) and 2005 (bottom panel). Values are averaged over the 3 summer months (JJA period) and individually over each summer month. Columns are given with a $0.5^\circ \times 0.5^\circ$ resolution for the JJA period and with a $1^\circ \times 1^\circ$ resolution for the monthly mean (see text for details).

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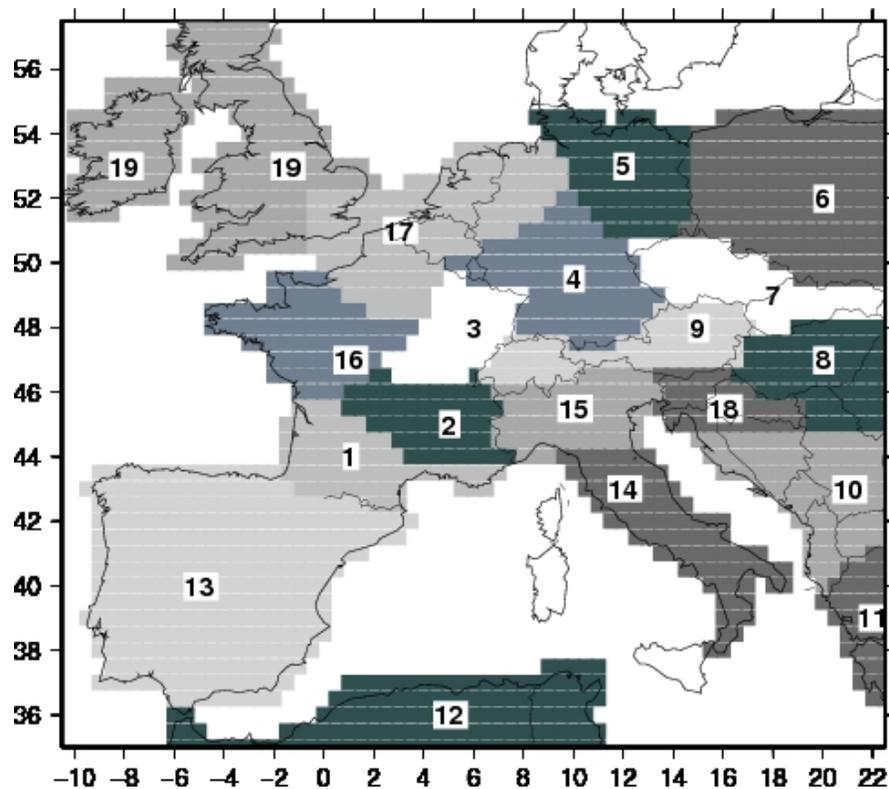


Fig. 7. Nineteen subdomains defined for SCIAMACHY and CHIMERE HCHO column comparison.

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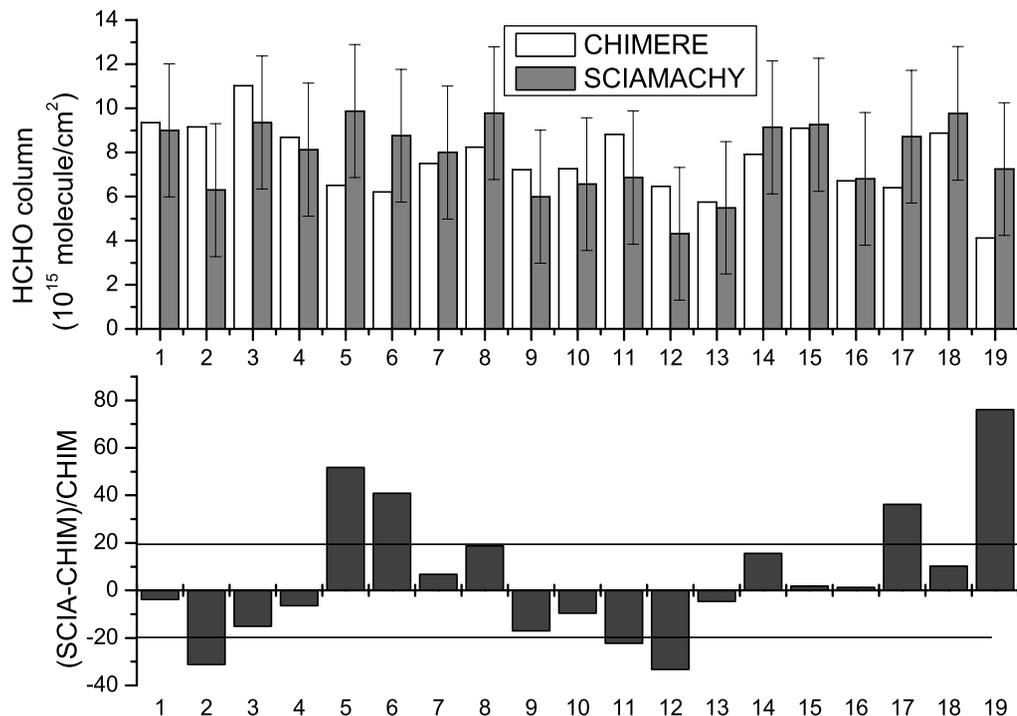


Fig. 8. Comparison of the mean HCHO tropospheric column measure by SCIAMACHY and simulated by CHIMERE in each subdomain (upper panel). The relative differences are given in the bottom panel.

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