

A global stratospheric bromine monoxide climatology based on the BASCOE chemical transport model

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Abstract

A new climatology of stratospheric BrO profiles based on dynamical and chemical indicators has been developed, with the aim to apply it to the retrieval of tropospheric BrO columns from space nadir measurements. The suitability of the adopted parameterization is evaluated based on three years of output data from the 3-D chemistry transport model BASCOE. The impact of the atmospheric dynamics on the stratospheric BrO distribution is treated by means of Br_y/ozone correlations build from 3-D-CTM model results, while photochemical effects are taken into account using stratospheric NO₂ columns as an indicator of the BrO/Br_y ratio. The model simulations have been optimized for bromine chemistry and budget, and validated through comparisons using an extensive data set of ground-based, balloon-borne and satellite limb (SCIAMACHY) stratospheric BrO observations.

1 Introduction

Inorganic bromine (Br_y=Br+BrO+BrONO₂+HOBr+HBr+BrCl+2Br₂) plays an important role in the stratosphere, as it is known to cause significant ozone destruction. Bromine monoxide (BrO) can significantly influence the chemical composition of the stratosphere and contribute to ozone depletion through catalytic reactions with NO₂, ClO, HO₂ and O (Lary, 1996). Although inorganic bromine is present in the stratosphere in small quantities relative to chlorine, its large ozone depletion efficiency makes stratospheric bromine contributing to mid-latitude ozone loss by about 25% and to polar ozone depletion, mainly through the BrO/ClO cycle (McElroy et al., 1986), up to 50%.

The origin of stratospheric inorganic bromine is both natural and anthropogenic. The major contribution to the stratospheric inorganic bromine budget comes from long-lived bromine-containing organic source gases, transported from the Earth's surface to the stratosphere, where they are converted into inorganic forms. The long-lived organic bromine species currently contribute to the stratospheric Br_y loading by

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about 16–17 pptv (Montzka et al., 2003). This estimate is based on a budget of observations of the main long-lived organic bromine source gases (CH_3Br , CBrClF_2 , CBrF_3 , $\text{CBrF}_2\text{CBrF}_2$). However, several recent studies based on measurements of stratospheric BrO by using remote-sensing UV-visible techniques from ground-based (Sinnhuber et al., 2002; Schofield et al., 2004, 2006; Hendrick et al., 2007; Theys et al., 2007), balloon-borne (Pundt et al., 2002; Salawitch et al., 2005; Dorf et al., 2006a, b, 2008) and space-borne limb (Sinnhuber et al., 2005; Sioris et al., 2006) instruments have inferred a total inorganic bromine loading of 18–25 pptv, suggesting that an additional contribution must be considered, possibly due to bromine release from short-lived biogenic organic compounds (such as CHBr_3 , CH_2Br_2 , CH_2BrCl , CHBr_2Cl , CHBrCl_2 , $\text{CH}_2\text{BrCH}_2\text{Br}$) or even through direct injection of inorganic bromine from tropospheric origin into the lower stratosphere (WMO, 2007). Enhanced Br_y due to very short-lived species (VSLs) can affect the ozone loss and chemical composition in the lower stratosphere and substantially influence ozone trend simulations (e.g. Salawitch et al., 2005; Feng et al., 2007).

Inorganic bromine compounds can also be present in the free-troposphere due to the decomposition of CH_3Br and VSLs bromine compounds. This is supported by observations from space (Wagner and Platt, 1998; Richter et al., 2002; Van Roozendaal et al., 2002), the ground (Hendrick et al., 2007; Theys et al., 2007) and balloon (Harder et al., 1998; Fitzenberger et al., 2000). Based on the observational estimates of bromine monoxide in the free-troposphere, modelling results (von Glasow et al., 2004; Lary, 2005; Yang et al., 2005) have shown that bromine might have a significant impact on tropospheric ozone (and on tropospheric chemistry in general), leading to a reduction in the O_3 mixing ratio of up to 40% locally. Further down in the boundary layer, inorganic bromine compounds have been shown to be responsible of complete ozone depletion events (ODEs) in the polar regions, during the so-called polar spring bromine explosion event (Hausmann and Platt, 1994; Kreher et al., 1997; Hönninger and Platt, 2002; Frieß et al., 2004; Simpson et al., 2007). Although in smaller quantities, tropospheric BrO emissions have also been identified over salt lakes (Hebestreit

et al., 1999), as well as in the marine boundary layer (Leser et al., 2003) and in volcanic plumes (Bobrowski et al., 2003).

Satellite UV-visible nadir instruments (such as GOME/ERS-2, SCIAMACHY/ENVISAT and GOME-2/MetOp-1) offer the unique capability to study and monitor BrO at the global scale (Chance, 1998; Richter et al., 2002; Van Roozendaal et al., 2002; Wagner et al., 1998). However, to quantitatively investigate and characterize the spatial and seasonal variations of tropospheric BrO, the measured total BrO vertical column densities (VCDs) must be resolved into their stratospheric and tropospheric contributions.

Here, we report on the development of a stratospheric BrO profile climatology designed for use in the retrieval of global tropospheric BrO VCDs from space-borne nadir observations. A new method for the estimation of the stratospheric BrO content is proposed, which is able to reproduce the important spatial and temporal variations of stratospheric BrO by using dynamical and chemical indicators. The climatology uses measured quantities to evaluate stratospheric BrO, which makes the approach well suited for satellite nadir retrieval since it guarantees that the sounded air masses and geophysical conditions are optimally represented. In Sect. 2, we briefly review the 3-D chemical transport model (CTM) which is at the heart of our study. We describe the set-up of bromine species implemented in the model, in the light of our current understanding of stratospheric bromine photochemistry and budget. The treatment of sulphate aerosols in the model is also presented, since it can have a substantial impact on the stratospheric chemistry. Comparisons between modelled results and correlative observations of O₃, NO₂ and BrO are shown in Sect. 3. The stratospheric BrO profile (and corresponding column) climatology is developed and presented in Sect. 4, together with an error analysis. Conclusions are given in Sect. 5.

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2 Model

The model used is the stratospheric 3-dimensional chemical transport model which forms the core of BASCOE, a 4-dimensional variational (4D-Var) chemical data assimilation system designed for the analysis and forecast of stratospheric ozone and chemical species (see <http://bascoe.oma.be>). The CTM is described in Errera and Fonteyn (2001), Daerden et al. (2007) and Errera et al. (2008). It includes a full description of stratospheric chemistry based on 57 chemical species and a parameterization of the microphysics of polar stratospheric clouds (PSCs). All chemical species are advected and interact through 144 gas-phase reactions, 48 photolysis reactions and 9 heterogeneous reactions, all listed in the Jet Propulsion Laboratory (JPL) compilation evaluation 14 (Sander et al., 2003). The CTM is driven by ECMWF operational analysis of temperatures and winds. The model is defined on 37 pressure levels from the surface to 0.1 hPa, on a 3.75° latitude and 5° longitude grid, with a model time step of 30 min. While the model extends down to the surface, it does not include any tropospheric process and is not expected to produce a realistic chemical composition below the tropopause.

The data used in this study result from a CTM run (no assimilation) initialized with analysis of MIPAS assimilation (details can be found in Errera et al., 2008). The simulations start on 1 April 2003 and end on 31 March 2006, covering three years of data.

2.1 Bromine species

The BASCOE 3-D CTM includes a detailed bromine chemistry scheme involving the photochemical reactions listed in Table 1. Inorganic bromine species interact through 30 gas-phase and heterogeneous photochemical reactions. Simulations include the reaction of $\text{BrONO}_2 + \text{O}(^3\text{P}) \rightarrow \text{BrO} + \text{NO}_3$, since several studies (Soller et al., 2001; Sinnhuber et al., 2002) found that this reaction can lead to increasing daytime BrO amount above 25 km, especially at low latitudes. Updated kinetic and photochemical data for bromine species are taken from the JPL 2006 compilation (Sander et al.,

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2006).

Our calculations assume a total inorganic bromine loading in the stratosphere of 23 pptv, in accordance with published results (Pfeilsticker et al., 2000; Salawitch et al., 2006; Hendrick et al., 2007, 2008a; Theys et al., 2007). The model includes the main tropospheric long-lived (LL) organic bromine sources: methyl bromide (CH_3Br), Halon-1211 (CBrClF_2) and Halon-1301 (CBrF_3). The volume mixing ratios (VMRs) have been scaled and made consistent with surface values measured in 1999 by Montzka et al. (2003); i.e. CH_3Br : 9.5 pptv, Ha-1211: 4.0 pptv and Ha-1301: 3.5 pptv (a constant 1 pptv of Br from Ha-2402 has been added to the measured 2.5 pptv of Ha-1301). The full conversion of LL organic bromine compounds into inorganic forms (occurring above ~ 25 km) leads to a stratospheric Br_y loading of 17 pptv, for air of 4–5 years mean age. To account for the bromine release from very short-lived species, an additional contribution of 6 pptv has been considered, consisting of 5 pptv of Br_y from dibromomethane (CH_2Br_2) and 1 pptv of tropospheric inorganic product gas directly injected at the tropopause. This method of implementing the short-lived bromine species is similar to what has been used in Feng et al. (2007).

A loss process of Br_y in the troposphere through washout due to the high solubility of HBr , has been implemented and is described by a washout time τ . This washout time is assumed to be constant throughout the troposphere and is fixed to 15 days. Model calculations often use a mean washout time τ of about 10 to 30 days (e.g. von Glasow et al., 2004; Sinnhuber et al., 2006).

2.2 Stratospheric aerosol settings

A particular attention has been paid to the treatment of the stratospheric sulphate aerosols and their impact on stratospheric bromine chemistry. The heterogeneous reactions on the surface of stratospheric aerosols can significantly influence the BrO/Br_y partitioning through its impact on the NO_x/NO_y ratio or directly through the heterogeneous reactions involving inorganic bromine species (Lary et al., 1996).

In the present work, an improved set-up of stratospheric aerosols has been imple-

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mented, derived from that used in Daerden et al. (2007). The aerosol surface area density is calculated based on the aerosol size distribution, which is assumed to follow a log-normal shape defined by the so-called RNS parameters (R : mean radius, N : total number density, S : standard deviation). The BASCOE aerosol configuration relies on a composite profile climatology of RNS parameter estimates for each month of the year and for 19 zonal bands. The RNS parameters have been estimated such that the corresponding $1\ \mu\text{m}$ aerosol extinction is consistent with SAGE-II extinction from Bingen et al. (2004) for latitudes between $\pm 60^\circ$ and with POAM III extinction from Fromm et al. (2003) for regions poleward of $\pm 60^\circ$. Adopting this approach for the treatment of stratospheric aerosols substantially improves the results. This is further verified in Sect. 3.1 where BASCOE NO_2 stratospheric columns are compared with observations from ground-based UV-visible instruments of the Network for the Detection of Atmospheric Composition Change (NDACC).

3 Verification of model results

Several studies have been conducted recently (Geer et al., 2006; Vigouroux et al., 2007; Daerden et al., 2007; Errera et al., 2008), where the BASCOE CTM has been extensively evaluated based on O_3 , NO_2 , H_2O , HNO_3 , N_2O and CH_4 observations. These results showed that BASCOE is able to reproduce important aspects of the stratospheric composition. Here, the reliability of the BASCOE results is further assessed for ozone, nitrogen dioxide (Sect. 3.1) and bromine monoxide (Sect. 3.2), by means of validation using ground-based, balloon and satellite limb observations.

3.1 Comparison of modelled and measured stratospheric O_3 and NO_2 columns

The measurement data were obtained from zenith sky UV-visible observations, performed at four stations which are part of the NDACC (<http://www.ndacc.org>). The observations are made during twilight periods for solar zenith angles (SZA) ranging

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from 86° to 91°. At all stations, the column densities of O₃ and NO₂ along the optical path (slant column) have been retrieved using the DOAS (Differential Optical Absorption Spectroscopy) technique (Platt and Stutz, 2006). The analysis is made for O₃ in the 435–555 nm wavelength interval and for NO₂ in the 420–470 nm range. Slant columns are converted into vertical columns using an optical path enhancement factor, or so-called air mass factor (AMF). Since ground-based measurements are made for high SZA, the retrieved O₃ and NO₂ columns are predominantly of stratospheric origin. For our comparison exercise, simulated ozone and nitrogen dioxide concentration profiles have been interpolated at the location corresponding to each station and for solar zenith angles of 90° (representative of the ground-based retrieved columns at sunrise and sunset), and then integrated between the tropopause and the highest model level.

Figure 1 presents the modelled and measured stratospheric O₃ columns for the period from April 2003 to March 2006 for Harestua (60.2° N, 10.7° E), Jungfraujoch (46.5° N, 7.9° E), Kerguelen (49.4° S, 70.3° E) and Rothera (67.6° S, 68.1° W). A good agreement is found for the entire period for Harestua, Jungfraujoch and Kerguelen with relative differences smaller than 15%. The observed seasonal and short-term dynamical changes are well captured by the simulations. At Rothera, the agreement is reasonable except during ozone-hole and summer conditions, where BASCOE overestimates the stratospheric O₃ columns by 50 to 100 DU. This confirms the findings of Errera et al. (2008). Note however that the measurements at Rothera have the tendency to underestimate the ozone columns by 10 to 15 DU for ozone hole conditions (Roscoe et al., 2001).

To a large extent, the lower stratospheric BrO concentration is controlled by NO₂ through the termolecular reaction $\text{BrO} + \text{NO}_2 + \text{M} \rightarrow \text{BrONO}_2$. Accurate modelling of nitrogen dioxide is thus a prerequisite to evaluate the stratospheric BrO amount.

Figure 2 shows the simulated and measured stratospheric NO₂ columns for Harestua, Jungfraujoch, Kerguelen and Rothera. The results are shown for sunset, but similar agreement is obtained for sunrise conditions. Most significant features of the observed seasonal and short-term NO₂ variabilities (related to photochemical and/or

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dynamical effects) are consistently reproduced by the simulations. However, the model shows a small tendency to underestimate the retrieved NO₂ columns. This might be related to the limited time/horizontal resolution and to the fact that the modelled columns are estimated at twilight, conditions where stratospheric NO₂ undergoes a rapid variation with time. Sensitivity tests show that the estimation of the modelled stratospheric NO₂ columns can be affected by this effect, up to 0.7×10^{15} molec/cm², with the largest discrepancies in summer. The absolute accuracy of the retrieved stratospheric NO₂ columns at NDACC stations can also be limited by (1) uncertainties related to the NO₂ absorption cross-sections and their temperature dependence (Koike et al., 1997; Roscoe et al., 1999; Vandaele et al., 2005), (2) AMF uncertainties associated with seasonal, meridian and diurnal variations of the NO₂ profile shape (Lambert et al., 1999). The overall uncertainty on the retrieved stratospheric NO₂ columns is of about 10%. Note that the possibility of a significant contribution to the measured column by tropospheric NO₂ from pollution is unlikely since the selected stations are located in clean areas. Nonetheless, the BASCOE model results agree in a satisfactory way with the observations (differences smaller than 1.2×10^{15} molec/cm²), knowing the uncertainties on both modelled and measured stratospheric NO₂ columns.

3.2 Comparisons of modelled and measured stratospheric BrO

The consistency between the stratospheric BrO profiles, as simulated by the BASCOE model, and correlative data has been verified. For this purpose, the modelled BrO profiles have been interpolated at the location and at the solar zenith angle of each observation. It should be noted that few measurements of stratospheric bromine monoxide overlap with the period covered by the simulations. The measurements reported in this study performed from ground-based, balloon and satellite platforms, are listed in Table 2. These correlative data set sample latitudes extending from polar to tropical regions, with the advantage that the resulting comparisons can be made under a large variety of photochemical regimes.

The stratospheric BrO vertical profiles obtained by the Microwave Limb Sounder

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(MLS) on the Aura satellite (Kovalenko et al., 2008) has not been used for our comparisons, since these measurements are sensitive to the BrO profile over a pressure range from 10 to 3.2 hPa, where only a small fraction of the total amount of stratospheric BrO is present. Furthermore, the estimated precision and accuracy of the data make them unsuitable for our comparison study.

3.2.1 Comparison to ground-based stratospheric BrO data

The ground-based observations have been conducted by IASB-BIRA at three NDACC stations: Harestua, Observatoire de Haute-Provence (OHP) and Reunion Island. The period covered by the simulations and the observations is April 2003–March 2006 at Harestua, February 2005–March 2006 at OHP and July 2004–July 2005 at Reunion Island. The measurements were obtained by zenith-sky UV-visible spectroscopy. The instrumental set-ups are described by Hendrick et al. (2007, 2008b) and Theys et al. (2007). The measured radiances are analysed using the DOAS technique. BrO differential slant columns densities (DSCDs) are retrieved in the 345–359 nm wavelength range, taking into account the spectral signature of BrO, NO₂, O₃, O₄, HCHO, OCIO and the Ring effect. The BrO cross-sections used are taken from Wilmouth et al. (1999). A detailed description of the BrO DOAS settings can be found in Theys et al. (2007) and Aliwell et al. (2002).

At Harestua and OHP, stratospheric BrO profiles are retrieved by applying a profiling technique to the measured DSCDs. The retrieval algorithm is based on the Optimal Estimation Method (Rodgers, 2000) and has been extensively described in Hendrick et al. (2007). It is based on a forward model that includes a stacked box photochemical model allowing to reproduce the strong diurnal variation of BrO at twilight. The stratospheric BrO profiles are retrieved for a solar zenith angle of 80°, separately for sunrise and sunset measurements. However, the retrieved BrO profiles can be converted to any SZA by using the photochemical model. The characterization of the information content of the retrieval shows that high sensitivity to the stratospheric BrO profile is obtained between 12- and 28-km altitude, with a vertical resolution of approximately

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8 km.

Ground-based BrO partial columns are calculated between 12–20 km and 20–28 km, in order to take benefit from the vertical resolution offered by the profiling technique, and are compared to BASCOE data. Figures 3 and 4 present the comparisons between modelled and ground-based BrO partial columns, at Harestua and OHP respectively. The error bars associated with the ground-based partial columns are estimates of the random and systematic uncertainties of the retrieval. The model and ground-based data are given at 80° SZA for sunset. At both stations, the modelled partial columns agree within 20–25% with the values derived from the ground-based observations. In particular, the seasonality of BrO, directly linked to the NO₂ seasonal cycle, is consistently captured by both model and ground-based data. Short-term variations linked to dynamical changes are also reproduced by both datasets in the lower stratosphere. Although not reproduced here, similar results are found with morning data.

At Reunion Island, the inversion of stratospheric BrO profiles by applying the profiling technique has been found to be difficult for two reasons: (1) the poorer performance of the instrument in terms of the signal to noise ratio, and (2) the smaller BrO absorption at tropical latitudes. Consequently, only stratospheric BrO columns have been retrieved, using an inversion method where the measured slant columns are fitted to radiative transfer simulations of the BrO slant column (see details in Theys et al., 2007). The error estimate associated to the retrieved stratospheric columns is slightly larger for Reunion-Island than for Harestua and OHP data. For this study, we compare the retrieved and modelled stratospheric BrO columns only at a single reference SZA (fixed at 80° for sunset). The comparison between modelled and ground-based stratospheric BrO columns is displayed in Fig. 5. The modelled BrO columns are calculated by integrating the stratospheric BrO profiles from the tropopause height, based on temperature profiles from ECMWF data, to the upper model level. Figure 5 shows that the modelled BrO columns are slightly smaller than the columns retrieved from ground-based measurements, but are within the error bars.

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3.2.2 Comparison with LPMA/DOAS balloon profiles

A detailed description of the LPMA/DOAS (Limb Profile Monitor of the Atmosphere/Differential Optical Absorption Spectroscopy) BrO measurements used here can be found in Harder et al. (1998, 2000), Pfeilsticker et al. (2000) and Dorf et al. (2006a, b, 2008). The LPMA/DOAS measurements use the solar occultation technique. The BrO profiles are retrieved using an Optimal Estimation Technique (Rodgers, 2000) applied to the BrO slant column densities measured during the ascent of the balloon. The accuracy of the technique is generally very good (lower than $\pm 12\%$, Harder et al., 1998) and the vertical resolution of the inverted stratospheric BrO profile is about 2 km. A further independent constraint on BrO, and hence on total stratospheric Br_y, is provided by Langley type observations of BrO absorption as function of total air mass above float altitude (>32 km) during each of balloon measurements (e.g., Dorf et al. 2006a).

For the present comparisons, three balloon flights originated from Kiruna, Aire sur l'Adour and Teresina are taken into account (see Table 2 for details). The BASCOE BrO profiles have been interpolated to the location of the balloon and at the SZA corresponding to the observation (varying during the ascent of the balloon) in order to have identical photochemical conditions between the balloon and the model profiles. Figure 6 shows the measured BrO balloon profiles together with the model calculation for the three flights. In general, we find a good agreement between the LPMA/DOAS and modelled BrO profiles. The observed differences are within the error bars of the balloon-borne profiles. These comparison results are similar to the ones presented in Feng et al. (2007) and Dorf et al. (2006a, b, 2008).

3.2.3 Comparison to SCIAMACHY limb profiles

The SCIAMACHY (Scanning Imaging Absorption Spectrometer for Atmospheric CHarotography) instrument is in operation on the ENVISAT platform since July 2002. In the nadir and limb viewing geometries, the SCIAMACHY instrument measures the sun-

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light scattered by the Earth's atmosphere or reflected by the surface whereas in the occultation mode, the direct solar or lunar light transmitted through the atmosphere is observed. The measurements are performed in eight spectral channels covering the 240–2400 nm wavelength range with a spectral resolution of 0.2 to 1.5 nm. A detailed description of the instrument and its characteristics can be found in Bovensmann et al. (1999). Among other geophysical data products, stratospheric BrO profiles can be retrieved from the SCIAMACHY limb spectral measurements (Rozanov et al., 2005; Sinnhuber et al., 2005; Sheode et al., 2006; Sioris et al. 2006) on the global scale (since ENVISAT is in polar sun-synchronous orbit). In this study, we use the BrO profiles derived at the Institute of Environmental Physics (IUP) at the University of Bremen (scientific product: version 3.2). A detailed description of the retrieval algorithm as well as the results of the error analysis and sensitivity studies can be found in (Rozanov et al., 2005) where, however, a precursor version of the retrieval processor (version 1.x) is discussed. The most recent retrieval version 3.2, used in this study, employs the same inversion algorithm as earlier versions (e.g., version 1.x described in the above cited paper) differing, however, quite strongly in the retrieval parameter settings. For example, a slightly different spectral range (338.0–356.2 nm) and a higher reference tangent height (about 35 km) were used, Levenberg-Marquardt iterative scheme was replaced by more common Newton-type iterations, and the regularization parameters were optimized. Furthermore, additional information on pressure and temperature provided by the European Centre for Medium-Range Weather Forecasts (ECMWF) were used. A more extensive description of the retrieval processor of version 3.2 is presented in (Hendrick et al., 2008b), where also the accuracy of the retrieved BrO profiles has been assessed. A detailed list of the retrieval parameter settings for the current and previous versions of the retrieval software can be found at the data archive web page of the IUP-Bremen (<http://www.iup.physik.uni-bremen.de/scia-arc>). SCIAMACHY limb BrO profiles are retrieved with the highest sensitivity in the 15–27 km altitude region and the vertical resolution is between 3 and 5 km (Rozanov et al., 2005; Hendrick et al., 2008b).

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For our comparison exercise, we have selected 11 complete days of data (the 15th of each month from April 2003 to February 2004). For each day, modelled and measured BrO partial columns have been calculated and compared for three altitude regions (15 to 19 km, 19 to 23 km and 23 to 27 km). Figure 7 shows the comparison results (zonal means, for SZA smaller than 85°), for the 15th of April, July, October of 2003 and January 2004. The error bars represent the standard deviation of the SCIAMACHY partial columns within each latitudinal bin plus an estimated systematic error of 10% (added in quadrature). The minimum error is determined by the detection limit of the SCIAMACHY limb measurements which is estimated to be about 10¹² molec/cm². Considering the uncertainties on satellite measurements, the agreement obtained is satisfying at all latitudes, for all altitude layers.

Comparison results for the 15–27 km BrO partial columns for different latitude bands (60° S–50° S, 40° S–30° S, 30° N–40° N and 50° N–60° N) are depicted in Fig. 8, for the 11 days of data (designated by their respective months on the x-axis). It shows that the agreement between BASCOE and SCIAMACHY partial columns is reasonably good; BASCOE data are most of the time within the error bars associated to the SCIAMACHY partial columns. The seasonality of BrO is also consistently captured at all latitudes by both SCIAMACHY and BASCOE data.

3.2.4 Discussion

As shown in previous sections, the model calculation reproduces the observed BrO generally very well, taking into account the overall errors of the measurements (about 20–25%), for the three different types of observations (ground-based, balloon, spaceborne limb measurements) used in this study. The seasonal and latitudinal variations are well captured by the model.

Model and measured data agree well assuming a total inorganic bromine loading of 23 pptv. The uncertainty on the contribution of VSLS to stratospheric Br_y has recently received a lot of attention, and is a topic stressed in the WMO report (2007). Furthermore, the quantification of the inorganic bromine amount at the lower stratospheric

entry levels is difficult, because it is controlled by processes with rather large uncertainties (surface emissions, transport and chemical lifetime of brominated short-lived species). A simple way to implement the bromine short-lived species in the model has been considered by assuming that the total Br_y is coming from a single effective source of CH_3Br (with a mixing ratio of 23 pptv at the surface). The stratospheric BrO results, from this test scenario, are displayed with green lines in Fig. 9, presenting the lower stratospheric BrO partial columns (12–20 km altitude range) at Harestua and OHP together with the measured data, for the period from January 2005 to March 2006. It can be seen that, in general, the stratospheric BrO modelled in this way is substantially smaller than the measured stratospheric BrO. Taking into account that the short-lived source gases are decomposed faster than methyl bromide, the set-up of short-lived species in the model as proposed in Sect. 2.1 leads to a better consistency between simulated and measured stratospheric BrO. This confirms, on a larger dataset, the findings of other studies (Feng et al., 2007; Dorf et al., 2006b, 2008), who found that the agreement between model and balloon data is significantly improved when having explicitly the short-lived source gases.

Sensitivity tests have been carried out in order to estimate the impact on the model results of the uncertainties on the main reaction rate constants (given in the JPL 2006 compilation), following the work presented in Sinnhuber et al. (2002) and Hendrick et al. (2008a). This gives a total error on the daytime BrO concentration smaller than 15%. Further photochemical uncertainties are kept small by considering in our comparisons, only data for SZA lower than 85° .

Moreover, stratospheric BrO is strongly dependent on NO_2 and the reasonable agreement between modelled and measured NO_2 columns at the NDACC stations (see Sect. 3.1), reinforce our confidence in the BrO simulations and in the fact that discrepancies due to non-appropriate simulation of NO_2 are rather small. Additional comparisons of BASCOE and SCIAMACHY limb profiles were performed for O_3 and NO_2 (details on SCIAMACHY limb retrieval can be found in Bracher et al., 2007 and Butz et al., 2006). A satisfactory agreement has been found for both species, in accor-

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dance with the conclusions of Sect. 3.1. These results further consolidate our modelled BrO results.

4 Stratospheric BrO climatology

4.1 General approach

5 The primary motivation behind the development of a new BrO climatology, is the retrieval of tropospheric BrO columns from satellite nadir observations, using a residual approach. To achieve this goal, the stratospheric BrO climatology has to meet specific requirements:

– Global coverage.

10 – The diurnal variation of BrO has to be taken into account, since the instruments are sounding the atmosphere under a large range of possible solar zenith angles. However, we restrict ourselves to SZA lower than 80° , because the retrieval of tropospheric BrO columns from space nadir observations at high solar zenith angles, is largely complicated due to two main reasons: (1) the measured signal is dominated by the absorption in the stratosphere, due to the very large photon paths in the stratosphere and the reduced sensitivity to the troposphere, and (2) stratospheric BrO has a sharp photochemical variation at twilight, that can lead to photochemical gradients along the slant stratospheric photon path, and even horizontal inhomogeneity of the BrO field within the measured pixel.

15 – Stratospheric BrO is highly variable in time and space, and depends on several parameters and atmospheric conditions. The BrO climatology must be able to reproduce with enough precision the BrO profiles for the large variety of possible scenarios. Misrepresentation of important patterns of stratospheric BrO (e.g. due to the effect of atmospheric dynamics) will lead to artefacts on the retrieved tropospheric BrO columns.

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- The stratospheric BrO climatology must provide an effective profile representative of the sounded atmosphere (averaged profile over the measured pixel).

The stratospheric BrO climatology proposed here is based on a parameterization which reflects the main dynamical and photochemical processes controlling the distribution of BrO in the stratosphere.

The stratospheric bromine monoxide profile can be written:

$$\text{BrO} = \text{Br}_y \times \frac{\text{BrO}}{\text{Br}_y} \quad (1)$$

Where $\text{Br}_y = \text{Br} + \text{BrO} + \text{BrONO}_2 + \text{HOBr} + \text{HBr} + \text{BrCl} + 2\text{Br}_2$ is the inorganic bromine profile accounting for all inorganic bromine species (active bromine and bromine reservoirs). Since inorganic bromine can be considered as a good chemical tracer due to its long chemical lifetime, Eq. (1) effectively separates the effects on the BrO vertical distribution due to the dynamics of the atmosphere and the photochemistry. For a given stratospheric air parcel (with a constant Br_y volume mixing ratio), rapid photochemical reactions between the various inorganic bromine species are taking place and affect the partitioning of BrO into the inorganic bromine species family (BrO/Br_y). In this study, these two different aspects are treated separately by developing two distinct climatologies. The development of a Br_y profile climatology is the focus of Sect. 4.2, while a partitioning profile climatology (BrO/Br_y) is presented in Sect. 4.3. The key requirements to derive a suitable stratospheric BrO profile from Eq. (1) is to obtain sufficient information about the dynamical and photochemical state of the sounded atmosphere.

4.2 Dynamics of the stratosphere

Since most of the inorganic bromine molecules are located below 35 km, we mainly focus on the lower and middle stratosphere. Hereafter, we will refer to the lower, middle and upper stratosphere to designate the stratospheric layers from the tropopause to 25 km, from 25 km to 35 km and above 35 km, respectively. As already mentioned,

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the distribution of inorganic bromine in the stratosphere results from the influence of dynamical processes. For typical time scales of the order of months, the transport of air masses in the stratosphere is governed by the mean meridional circulation, which brings inorganic bromine produced in the tropical lower and middle stratosphere to extra-tropical latitudes. In principle, a monthly zonally-averaged Br_y profile climatology is able to reproduce in a satisfactory way the relevant spatial and temporal variations of stratospheric inorganic bromine due to the meridional circulation (e.g. ascending and descending advection in the tropical and polar regions respectively, Holton, 1995). However it fails to represent the Br_y variability due to shorter dynamical processes. A more sophisticated climatology is thus needed to reproduce the effects on Br_y due to short-term transport. At the scale of several days, the motion of stratospheric air parcels is driven by zonal winds. As a first approximation, it can be assumed that the air masses are transported adiabatically (motion along quasi-horizontal isentropic surfaces). Zonal winds tend to eliminate the longitudinal gradients of tracers of the stratospheric dynamics (as Br_y). The effect on the distribution of Br_y varies as a function of altitude, latitude and season. Such short-term dynamical effects are well accounted for by the BASCOE 3-D CTM since it uses 6 hourly ECMWF operational analysis of winds and temperatures. This is demonstrated by the fact that short-term changes in measured O_3 columns are consistently captured by the model (see Fig. 1 in Sect. 2.1). The important concept introduced in this work to reproduce the effect on stratospheric Br_y due to short-term transport, is to use stratospheric ozone as a proxy for the dynamical state of the stratosphere. For non-chlorine activated conditions, ozone in the lower and middle stratosphere can be considered as a tracer of the dynamics of the stratosphere (chemical lifetime about one to several months). Since stratospheric inorganic bromine and ozone are both produced mainly in the tropical lower stratosphere and are sensitive to the stratospheric dynamics in a similar way, a correlation between Br_y and O_3 is expected. This is illustrated in Fig. 10, where the O_3 and Br_y volume mixing ratios, as simulated by the BASCOE model, are plotted for a low stratospheric isentropic level for the 1 March 2004 in the Northern Hemisphere. The large dynamical patterns, charac-

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teristic of this season, are obviously reproduced by both ozone and Br_y VMRs. Based on this noticeable property, we propose a stratospheric Br_y concentration profile climatology based on a simple parameterization using several inputs: month, latitude and stratospheric ozone column. The dependency of the Br_y climatology with the month of the year and the latitude allows to reproduce the effect on Br_y due to the seasonal and latitudinal variations of the meridional circulation, and implicitly for the changes in the zonal winds.

The inorganic bromine profile climatology has been generated by considering the set of Br_y profiles, simulated by BASCOE for the period from April 2003 to March 2006. The Br_y profiles have been interpolated on a regular altitude grid (step: 1 km) covering the stratosphere and classified by month, latitude (18 bands defined by the following boundaries: -90°, -80°, . . . , 80°, 90°) and O₃ column (19 blocks of 25 Dobson unit (DU) width centred around 125, 150, . . . , 550, 575 DU). The inorganic bromine profile climatology is build by calculating the average of the Br_y profiles corresponding to each month, latitude band, and O₃ column bin. Different weights have been assigned to the Br_y profiles to account for the fact that the model cells vary in size with latitude and thus must have different contributions to the mean (this effect is particularly important close to the poles).

In order to gain better insight into the consistency of the proposed Br_y parameterization, Fig. 11 illustrates the correlation between stratospheric O₃ and Br_y vertical columns, as simulated by BASCOE, for two particular cases: (1) northern mid-latitudes (40° N–50° N) in March, (2) southern high-latitudes (70° S–60° S) in September. For northern mid-latitudes, the correlation between O₃ and Br_y vertical columns is excellent. The correlation coefficient is equal to 0.98. For southern high-latitudes in spring, the correlation is deteriorated since stratospheric ozone can no longer be considered as a dynamical tracer (due to its destruction by chlorine species, activated through heterogeneous reactions on the surface of PSCs), and because of the presence of the strong polar dynamical vortex. However, it can be seen that there is only a moderate dispersion of Br_y columns around the mean, for a given O₃ column interval. The rele-

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vance of the proposed Br_y climatology will be assessed and consolidated in Sect. 4.5, through a comprehensive error analysis.

As already mentioned in Sect. 4.1, the bromine monoxide profile climatology is intended to be used for the retrieval of tropospheric BrO columns from satellite nadir observations. In a first step of the calculation of BrO profiles (see Eq. 1), stratospheric Br_y concentration profiles can be derived by using the Br_y profile parameterization established here, together with the O₃ columns retrieved from the observations. This has several advantages: (1) it is simple and requires only limited computation resources, (2) the ozone column is a standard product which is retrieved operationally with an excellent accuracy (error lower than 1 or 2%), and (3) it guarantees to represent the same air masses and geophysical conditions as the BrO retrieval. However, the retrieved total O₃ column contains a small contribution from tropospheric ozone. This can not be accounted for in the Br_y climatology, since BASCOE does not include any tropospheric processes and is not expected to produce realistic results below the tropopause. In practice, this can be (at least partly) corrected for by subtracting, from the retrieved total O₃ column, a tropospheric contribution estimated from a tropospheric ozone climatology (e.g. Logan, 1999).

4.3 Bromine monoxide photochemistry

As mentioned earlier, the bromine monoxide profile is obtained by Eq. (1) through a partitioning factor (BrO/Br_y) calculated as a function of the altitude, and applied to the inorganic bromine profile. In order to evaluate the ratio (BrO/Br_y), it is necessary to identify the dominant photochemical regimes affecting the bromine partitioning during daytime in the lower and middle stratosphere. A detailed description of the gas-phase and heterogeneous stratospheric bromine chemistry can be found in Lary (1996) and Lary et al. (1996); see also a list of the most relevant reactions in Table 1.

For unperturbed atmospheric conditions, bromine monoxide is in photochemical equilibrium with the bromine reservoirs (mostly BrONO₂ and HOBr) formed by the re-

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The loss processes of BrONO_2 and HOBr are dominated by the photolysis and the reaction with O . Note that bromine nitrate can be converted by heterogeneous reactions on sulphate aerosols into HOBr :



This reaction is expected to play an important role on the partitioning factor before sunrise, but a smaller one on daytime bromine photochemistry (for SZA lower than 80°), because of the efficient photolysis of HOBr .

Hence during daytime, bromine nitrate is the major bromine reservoir. As a first approximation, the bromine partitioning factor can be expressed as follows:

$$\frac{\text{BrO}}{\text{Br}_y} \approx \frac{\text{BrO}}{\text{BrO} + \text{BrONO}_2} = \frac{1}{1 + \frac{K_{\text{BrONO}_2} \text{NO}_2}{L_{\text{BrONO}_2}}} \quad (2)$$

Where K_{BrONO_2} and L_{BrONO_2} are the formation and loss rates of BrONO_2 , respectively; NO_2 is the nitrogen dioxide concentration. This relation expresses the balance between the production and the loss of BrONO_2 , which affects directly the bromine partitioning factor (BrO/Br_y). It has to be emphasized that (1) the reaction rate of formation of BrONO_2 varies with altitude through a temperature dependence, and (2) the loss rate of BrONO_2 (mostly photolysis) varies strongly with altitude and solar zenith angle. A sensitivity test shows that it depends only slightly on the O_3 column. Consequently, the partitioning ratio BrO/Br_y responds very rapidly to any change in SZA and/or NO_2 concentration.

We propose a partitioning factor BrO/Br_y profile climatology based on a classification according to: month, latitude, stratospheric NO_2 column and solar zenith angle. For

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this reason, we have calculated the modelled BrO/Br_y profiles for the complete period (April 2003 to March 2006), on the altitude grid already used for the Br_y climatology. We have selected the BASCOE data corresponding to morning conditions, since the present climatology is designed to be applied for satellite instruments having morning overpasses (GOME, SCIAMACHY, GOME-2). The partitioning factor profile climatology is build by averaging the BrO/Br_y profiles on grids of SZA (5° bins, SZA lower than 80°) and stratospheric NO₂ columns (25 blocks of 0.25×10¹⁵ molec/cm² width centred around 0.125, 0.375, . . . , 5.875, 6.125×10¹⁵ molec/cm²).

In order to illustrate the concept of the partitioning (BrO/Br_y) climatology, Fig. 12 displays the ratio between the stratospheric BrO and Br_y columns as a function of the stratospheric NO₂ columns, simulated by BASCOE for August and December and for two solar zenith angles ranges (SZA≤60° and 60°<SZA≤80°).

The inverse dependency of BrO/Br_y with NO₂ is obvious, as well as the diurnal variation of BrO. However a striking feature can be observed in Fig. 12c, where two photochemical regimes are present in August for low stratospheric NO₂ columns. The lower branch of the scatter plot corresponds to perturbed atmospheric conditions in the Antarctic polar vortex (close to the terminator). Indeed, the polar stratospheric clouds (formed during the winter polar night) provide the surface for a number of heterogeneous reactions leading to the denoxification process and the activation of chlorine species (see e.g. Solomon, 1999). As a consequence, the lower stratospheric inorganic bromine is almost exclusively partitioned between BrO and BrCl (formed by the reaction between bromine and chlorine monoxides), and the bromine photochemistry deviates from the standard regime (essentially controlled by stratospheric NO₂).

It should be noticed that if we restrict ourselves to unperturbed conditions, a parameterization of BrO/Br_y with additional entries as the latitude and month is relatively redundant. In first approximation, the stratospheric NO₂ column is mainly determined by day length (photolysis of the nitrogen reservoirs) and the solar zenith angle (affecting the diurnal equilibrium NO₂/NO). Nevertheless, a general partitioning factor climatology using a classification by latitude and month, is able to reproduce (to some extend)

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the different photochemical regimes, associated to low NO₂ columns, experienced by the stratospheric inorganic bromine species.

From Fig. 12, it can be seen that there is a moderate dispersion of the data around the mean curves. It suggests that the error made by using the adopted parameterization of BrO/Br_y is rather small. This will be established in Sect. 4.5 through an error assessment.

The advantage of using the stratospheric NO₂ column to parameterize the BrO/Br_y ratio is that, besides the simplicity of the method, this information is easily accessible as an intermediate product in the retrieval of total and tropospheric NO₂ columns from space nadir measurements (e.g. Boersma et al., 2004).

4.4 Long-term trend in stratospheric bromine

Due to the evolution of the organic bromine emissions (Montzka et al., 2003) and the impact of the Montreal Protocol, a positive trend in stratospheric bromine of about +2.5% per year has been observed for the 1995–2001 period, while a negative trend of about –1% per year has been identified between 2001 and 2005 (Hendrick et al., 2008a). Since model fields used in this study cover three years of data from April 2003 to March 2006, the stratospheric BrO climatology can be generalized to account for long-term trend, by adding the year as a new entry of the climatology, in the form of a correction factor.

4.5 Results and error analysis

In this section, the stratospheric BrO vertical columns estimated from the Br_y and BrO/Br_y climatologies are presented. The suitability of the approach described in the previous sections is addressed through an error analysis. To simplify the discussion, we restrict ourselves to the vertical columns at the time of GOME overpass (similar results are obtained for instruments with late morning or early afternoon overpasses, as SCIAMACHY, GOME-2 and OMI). The instrument, aboard the ESA/ERS-2 satellite,

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flies in a sun-synchronous orbit, implying that fixed latitudes are always sounded at the same local time. For the period from April 2003 to March 2006, the BASCOE model data at GOME overpass has been selected and the stratospheric BrO, Br_y, O₃ and NO₂ vertical columns have been calculated (tropopause heights estimated from ECMWF pressure and temperature profiles, are used). The climatological stratospheric profiles of Br_y concentration and partitioning ratio (BrO/Br_y) were then evaluated by applying the parameterization, respectively to the calculated O₃ and NO₂ vertical columns. The climatological stratospheric BrO profiles and columns are calculated by Eq. (1). For each month and latitude band, the mean of the BrO columns from the climatology is estimated, together with the root mean square of the difference between the parameterized and the BASCOE modelled stratospheric BrO columns. Results are presented in Fig. 13. The same approach has been adopted for the Br_y columns. In order to assess the importance of the errors originating from the parameterization of Br_y and BrO/Br_y, both contributions have been considered separately using the following equation (Br_y and BrO/Br_y are independent variables):

$$\sigma_{\text{BrO}}^2 = (\text{BrO}/\text{Br}_y)^2 \sigma_{\text{Br}_y}^2 + (\text{Br}_y)^2 \sigma_{\text{BrO}/\text{Br}_y}^2 \quad (3)$$

Figure 13 shows the relative error on the stratospheric BrO columns due to Br_y and BrO/Br_y, respectively. The stratospheric Br_y vertical columns derived from the climatological profiles and the ratio between the BrO and Br_y stratospheric columns, are also presented.

From Fig. 13, it can be stated that the new stratospheric BrO climatology provides reliable results with a precision better than 14%. It also reveals that:

- the Br_y climatology produces higher Br_y columns at mid- and high-latitudes than for tropical regions, as expected from the dependency of the tropopause height with latitude. One can notice smaller values in the Antarctic polar spring, which are mainly related to the air subsidence in the polar vortex.
- except for perturbed chemistry conditions, the errors on the stratospheric BrO

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columns are dominated by the errors on Br_y . These errors are maximum at mid- and high-latitudes in winter and early spring during major changes in stratospheric dynamics, leading to a larger scatter in the Br_y /ozone correlations. The moderate errors on Br_y during Antarctic ozone hole conditions indicate that Br_y shows small variations within the vortex since the air over the polar region is well isolated. The use of the O_3 column helps to identify the vortex and discern air masses with Br_y amount substantially different.

- the BrO/Br_y factor exhibits a strong seasonal variation at mid- and high-latitudes, related to the seasonal cycle of stratospheric NO_2 .
- the relative error on the stratospheric BrO columns due to errors on the partitioning factor BrO/Br_y is remarkably small for usual photochemical regimes (lower than 8%). For perturbed chemistry conditions, the standard deviation can reach 12% of the BrO column due to the presence of various photochemical scenarios and dynamical mixing of air with different amount of NO_x .

An important aspect to emphasize is, that the BASCOE stratospheric O_3 and NO_2 columns used to build the stratospheric BrO climatology have been validated through comparisons with measured columns obtained from ground-based observations at mid- and high-latitudes in both hemispheres (see Figs. 1 and 2 in Sect. 3.1). The important observed changes in stratospheric O_3 and NO_2 are consistently reproduced by the model, except for polar ozone in the Southern Hemisphere in ozone-hole conditions, where BASCOE tends to overestimate the stratospheric O_3 columns. In practice, however, this is expected to have a marginal impact on the evaluation of stratospheric BrO profiles from our climatology since Br_y show small variations and correlates only slightly with stratospheric ozone for these conditions (see Fig. 11 in Sect. 4.2).

In summary, the error on the stratospheric BrO column evaluated by the BASCOE climatology, is limited by the accuracy of the modelled BrO. By considering the contribution of the main sources of errors added in quadrature, the overall uncertainty on stratospheric BrO can be kept below 30% under most observational conditions.

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5 Conclusions

We have presented a new global climatology of stratospheric bromine monoxide, primarily designed for the retrieval of tropospheric BrO columns from satellite nadir instruments (as GOME, SCIAMACHY and GOME-2), but also of interest for a wider range of applications. The impact of the stratospheric dynamics on the stratospheric BrO distribution is accounted for by a parameterization based on the ozone column, while the effect of photochemistry on stratospheric BrO is determined by considering the stratospheric NO₂ columns and the solar zenith angle. The adopted parameterization is evaluated based on three years of output data from the BASCOE chemical transport model. Model simulations include full gas phase chemistry and relevant heterogeneous reactions, while dynamics is driven by ECMWF wind fields. Regarding the bromine species, the model calculations use up-to-date photochemistry and a realistic total bromine budget (including a contribution from short-lived bromine source gases).

Although primarily based on model simulations, the climatology is also closely linked to an extensive set of stratospheric BrO observations from ground-based, balloon-borne and satellite (SCIAMACHY) platforms. The excellent overall agreement between the model and the observations demonstrates (1) the level of consistency of the currently available BrO observational data set, (2) the ability of our model to reproduce the main mechanisms controlling the distribution of stratospheric bromine monoxide. We estimate that the overall uncertainty of the resulting stratospheric BrO profiles is lower than 30%. The consequences to infer tropospheric BrO columns from satellite nadir measured total BrO columns and the present BASCOE stratospheric BrO climatology will be investigated in a future paper.

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part of the Network for the Detection of Atmospheric Composition Change (NDACC) and is publicly available (see <http://www.ndacc.org>).

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Table 1. Bromine reactions included in the BASCOE model.

| No. | Reaction | No. | Reaction |
|-------|---|-------|---|
| (R1) | $\text{BrO} + \text{NO}_2 + \text{M} \rightarrow \text{BrONO}_2 + \text{M}$ | (R21) | $\text{BrONO}_2 + \text{O} \rightarrow \text{BrO} + \text{NO}_3$ |
| (R2) | $\text{BrO} + \text{HO}_2 \rightarrow \text{HOBr} + \text{O}_2$ | (R22) | $\text{BrONO}_2 + \text{HCl} \xrightarrow{\text{Het}} \text{BrCl} + \text{HNO}_3$ |
| (R3) | $\text{BrONO}_2 + \text{H}_2\text{O} \xrightarrow{\text{Het}} \text{HOBr} + \text{HNO}_3$ | (R23) | $\text{HOBr} + \text{HCl} \xrightarrow{\text{Het}} \text{BrCl} + \text{H}_2\text{O}$ |
| (R4) | $\text{BrO} + \text{NO} \rightarrow \text{Br} + \text{NO}_2$ | (R24) | $\text{HOBr} + \text{HBr} \xrightarrow{\text{Het}} \text{Br}_2 + \text{H}_2\text{O}$ |
| (R5) | $\text{BrO} + \text{O} \rightarrow \text{Br} + \text{O}_2$ | (R25) | $\text{BrONO}_2 + h\nu \rightarrow \text{BrO} + \text{NO}_2$ |
| (R6) | $\text{BrO} + \text{OH} \rightarrow \text{HO}_2 + \text{Br}$ | (R26) | $\text{BrONO}_2 + h\nu \rightarrow \text{Br} + \text{NO}_3$ |
| (R7) | $\text{BrO} + \text{ClO} \rightarrow \text{Br} + \text{ClOO}$ | (R27) | $\text{HOBr} + h\nu \rightarrow \text{Br} + \text{OH}$ |
| (R8) | $\text{BrO} + \text{ClO} \rightarrow \text{Br} + \text{ClOO}$ | (R28) | $\text{BrCl} + h\nu \rightarrow \text{Br} + \text{Cl}$ |
| (R9) | $\text{BrO} + \text{ClO} \rightarrow \text{BrCl} + \text{O}_2$ | (R29) | $\text{BrO} + h\nu \rightarrow \text{Br} + \text{O}$ |
| (R10) | $\text{BrO} + \text{BrO} \rightarrow 2\text{Br} + \text{O}_2$ | (R30) | $\text{Br}_2 + h\nu \rightarrow 2\text{Br}$ |
| (R11) | $\text{BrO} + \text{BrO} \rightarrow \text{Br}_2 + \text{O}_2$ | (R31) | $\text{CH}_3\text{Br} + h\nu \rightarrow \text{Br} + \text{CH}_3$ |
| (R12) | $\text{Br} + \text{OCIO} \rightarrow \text{BrO} + \text{ClO}$ | (R32) | $\text{CH}_3\text{Br} + \text{O}(^1\text{D}) \rightarrow \text{Br} + \text{products}$ |
| (R13) | $\text{Br} + \text{HO}_2 \rightarrow \text{HBr} + \text{O}_2$ | (R33) | $\text{CH}_3\text{Br} + \text{OH} \rightarrow \text{Br} + \text{H}_2\text{O} + \text{products}$ |
| (R14) | $\text{Br} + \text{CH}_2\text{O} \rightarrow \text{HBr} + \text{HCO}$ | (R34) | $\text{CBrClF}_2 + h\nu \rightarrow \text{Br} + \text{Cl} + \text{products}$ |
| (R15) | $\text{Br} + \text{O}_3 \rightarrow \text{BrO} + \text{O}_2$ | (R35) | $\text{CBrClF}_2 + \text{O}(^1\text{D}) \rightarrow \text{Br} + \text{Cl} + \text{products}$ |
| (R16) | $\text{Br}_2 + \text{OH} \rightarrow \text{HOBr} + \text{Br}$ | (R36) | $\text{CBrF}_3 + h\nu \rightarrow \text{Br} + \text{products}$ |
| (R17) | $\text{HOBr} + \text{O} \rightarrow \text{BrO} + \text{OH}$ | (R37) | $\text{CBrF}_3 + \text{O}(^1\text{D}) \rightarrow \text{Br} + \text{products}$ |
| (R18) | $\text{HBr} + \text{O} \rightarrow \text{Br} + \text{OH}$ | (R38) | $\text{CH}_2\text{Br}_2 + h\nu \rightarrow 2\text{Br} + \text{products}$ |
| (R19) | $\text{HBr} + \text{OH} \rightarrow \text{Br} + \text{H}_2\text{O}$ | (R39) | $\text{CH}_2\text{Br}_2 + \text{OH} \rightarrow 2\text{Br} + \text{products}$ |
| (R20) | $\text{HBr} + \text{O}(^1\text{D}) \rightarrow \text{Br} + \text{OH}$ | | |

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Table 2. Summary of the BrO observation data sets used for the comparisons with model results.

| Observation type | Location | Period | Reference |
|--|-------------------|-------------------|-------------------------|
| Ground-based DOAS (IASB-BIRA) | | | |
| Harestua, Norway | 60.22° N 10.75° E | Apr 2003–Mar 2006 | Hendrick et al. (2007) |
| Observatoire de Haute-Provence, France | 43.94° N 5.71° E | Feb 2005–Mar 2006 | Hendrick et al. (2008b) |
| Reunion Island, France | 21.06° S 55.47° E | Jul 2004–Jul 2005 | Theys et al. (2007) |
| Balloon LPMA/DOAS (U. Heidelberg) | | | |
| Kiruna, Sweden | 67.84° N 21.06° E | 24 March 2004 | Dorf et al. (2006a) |
| Aire sur l'Adour, France | 43.70° N 0.25° W | 9 October 2003 | Dorf et al. (2006a) |
| Teresina, Brazil | 5.1° S 42.9° W | 17 June 2005 | Dorf et al. (2008) |
| SCIAMACHY limb (U. Bremen) | Global | Apr 2003–Feb 2004 | Rozanov et al. (2005) |

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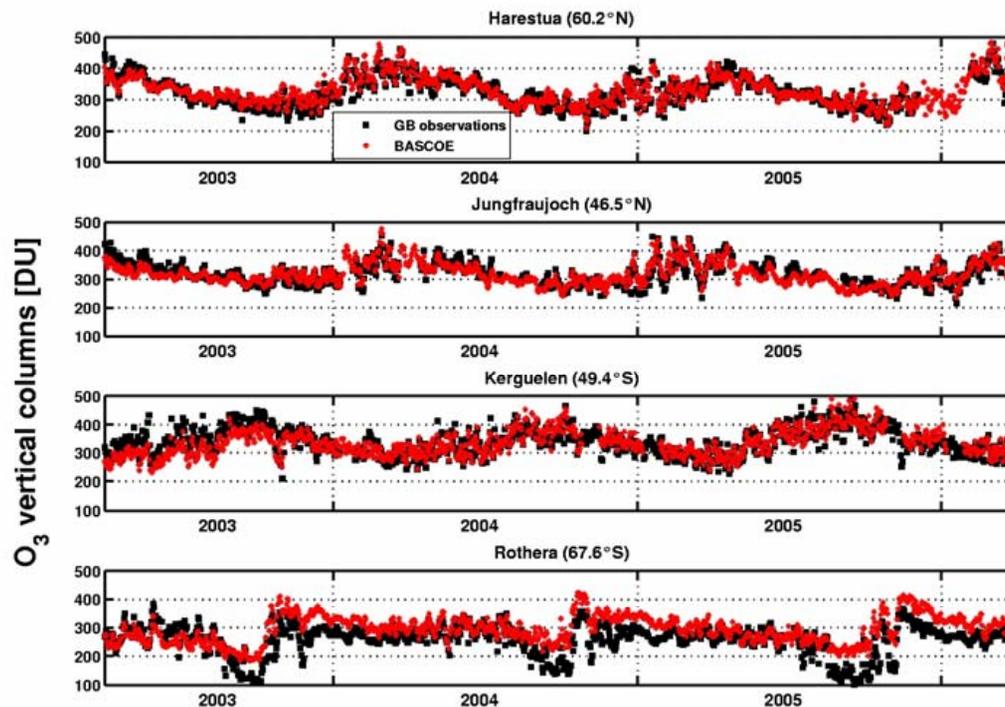


Fig. 1. Time-series of sunset stratospheric O₃ vertical columns from BASCOE analysis and ground-based UV-visible observations at 4 NDACC stations: Harestua, Jungfrauoch, Kerguelen and Rothera.

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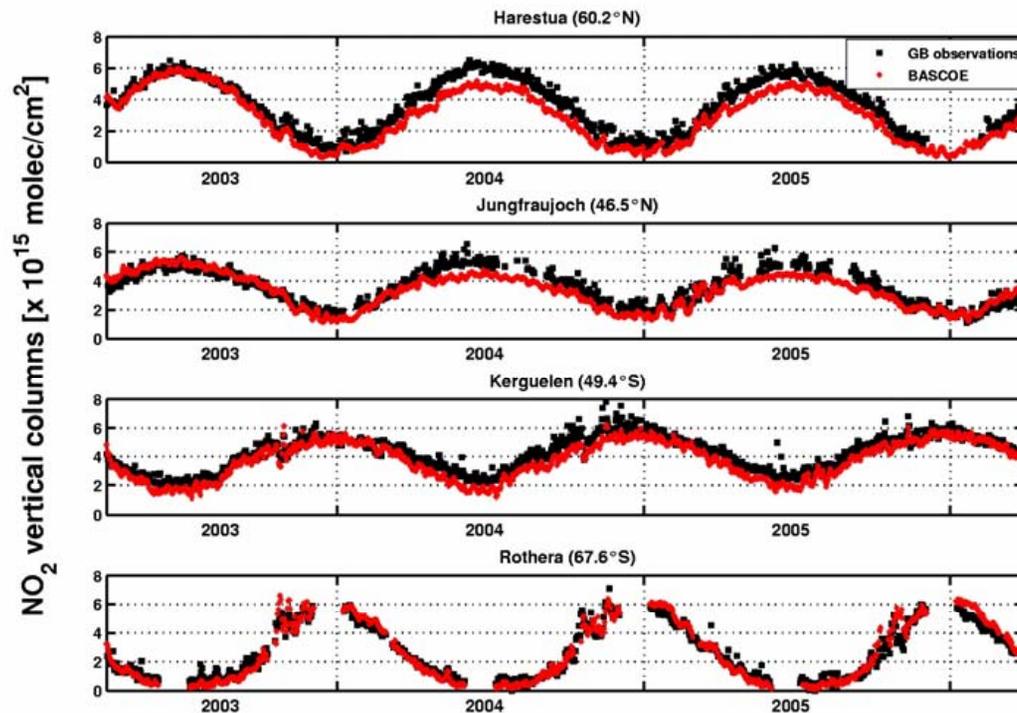


Fig. 2. Time-series of sunset stratospheric NO₂ vertical columns from BASCOE analysis and ground-based UV-visible observations at 4 NDACC stations: Harestua, Jungfraujoch, Kerguelen and Rothera.

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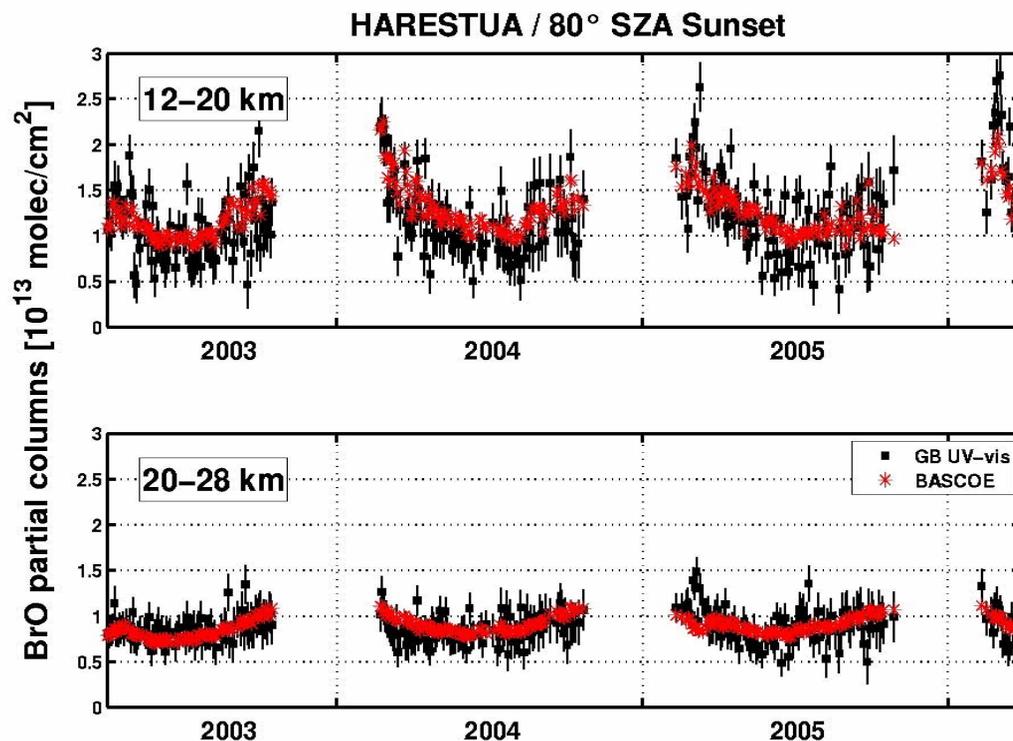


Fig. 3. Comparison between modelled and ground-based stratospheric BrO partial columns, at Harestua for 80° SZA sunset. The BrO profiles are integrated between 12–20 km (top) and 20–28 km (bottom).

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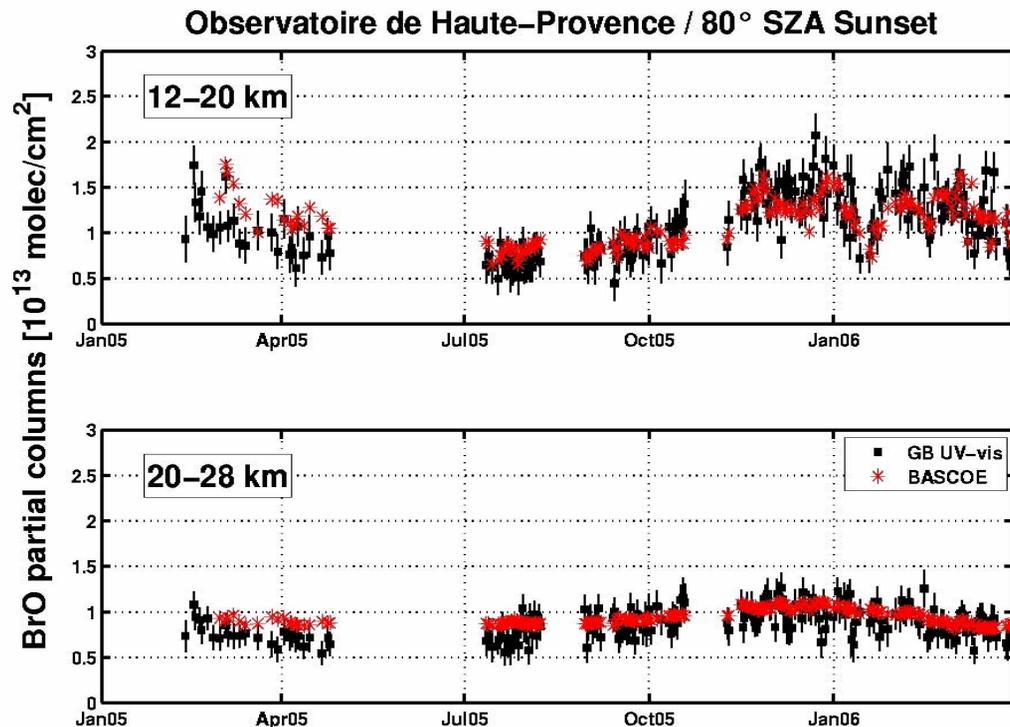


Fig. 4. Same as Fig. 3, at Observatoire de Haute-Provence.

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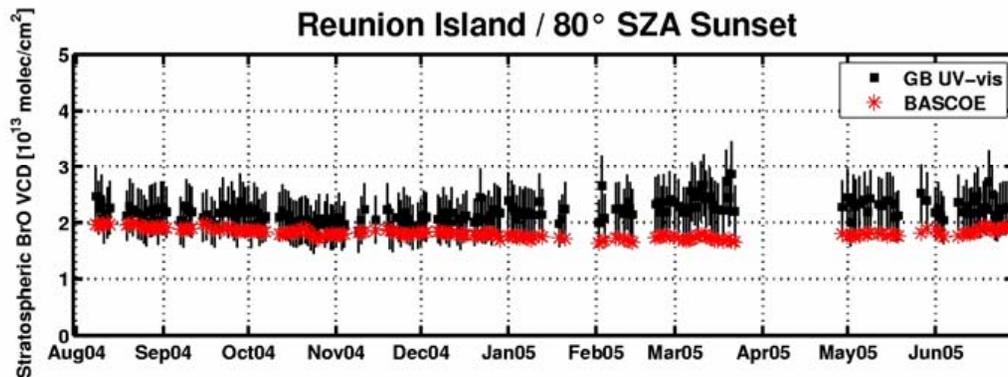


Fig. 5. Comparison between modelled and ground-based stratospheric BrO columns, at Reunion Island for 80° SZA sunset.

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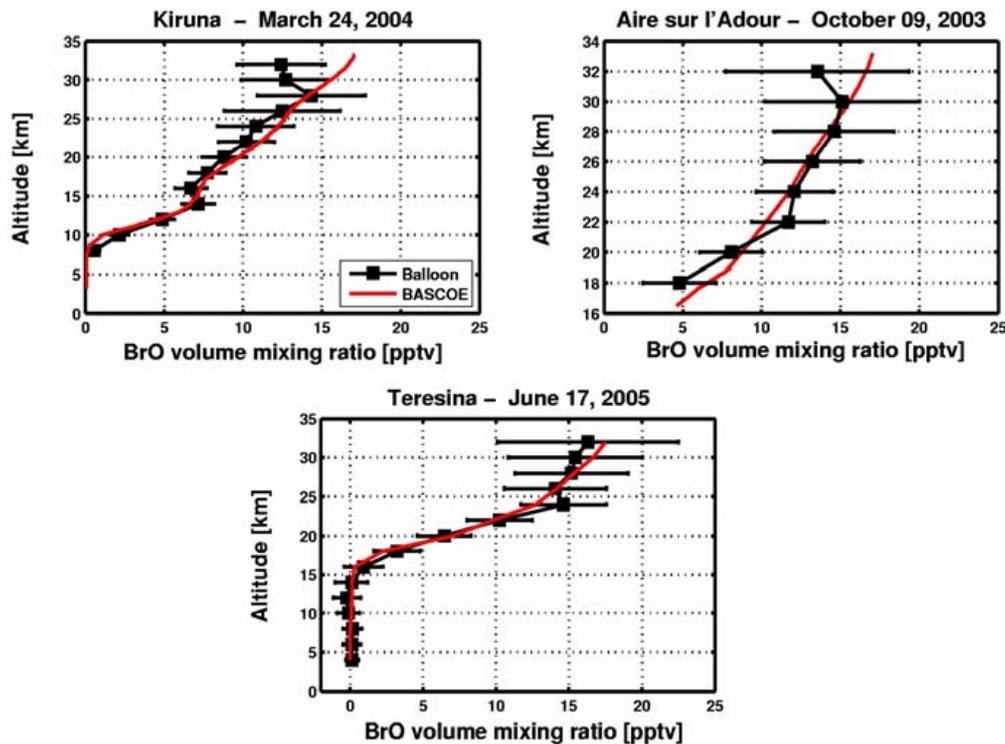


Fig. 6. Comparison between modelled and LPMA/DOAS ascent balloon BrO profiles at Kiruna, Aire sur l'Adour and Teresina.

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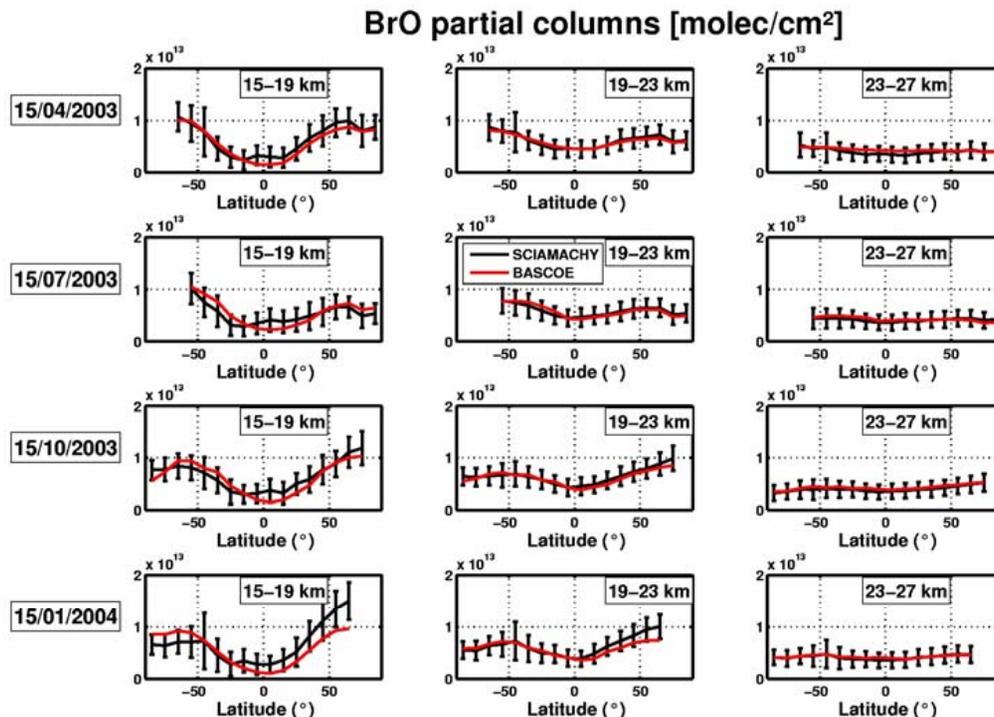


Fig. 7. Zonal means of BrO partial columns (SZA lower than 85°), for three altitude ranges (left plots: 15 to 19 km, centre plots: 19 to 23 km, right plots: 23 to 27 km) calculated from SCIAMACHY limb profiles and BASCOE modelled profiles for the 15 April, July, October 2003 and January 2004.

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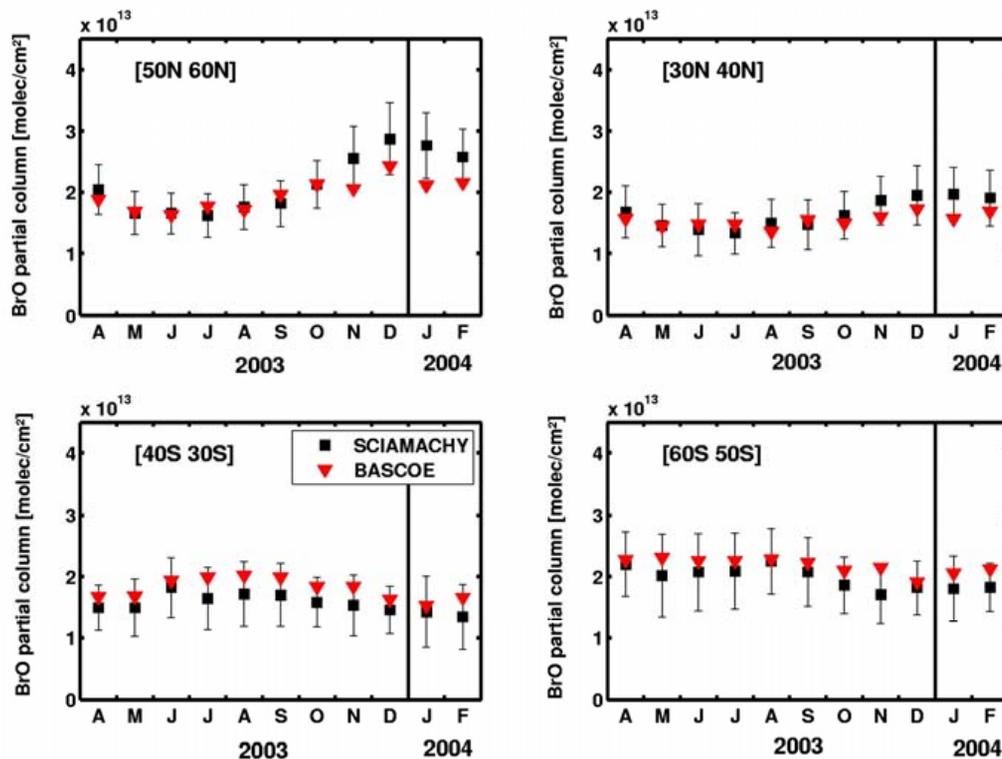


Fig. 8. Comparison of the 15–27 km BrO partial columns (zonal means) calculated from the SCIAMACHY limb and BASCOE profiles for 11 days of data sampling the period from April 2003 to February 2004 and for different latitude bands.

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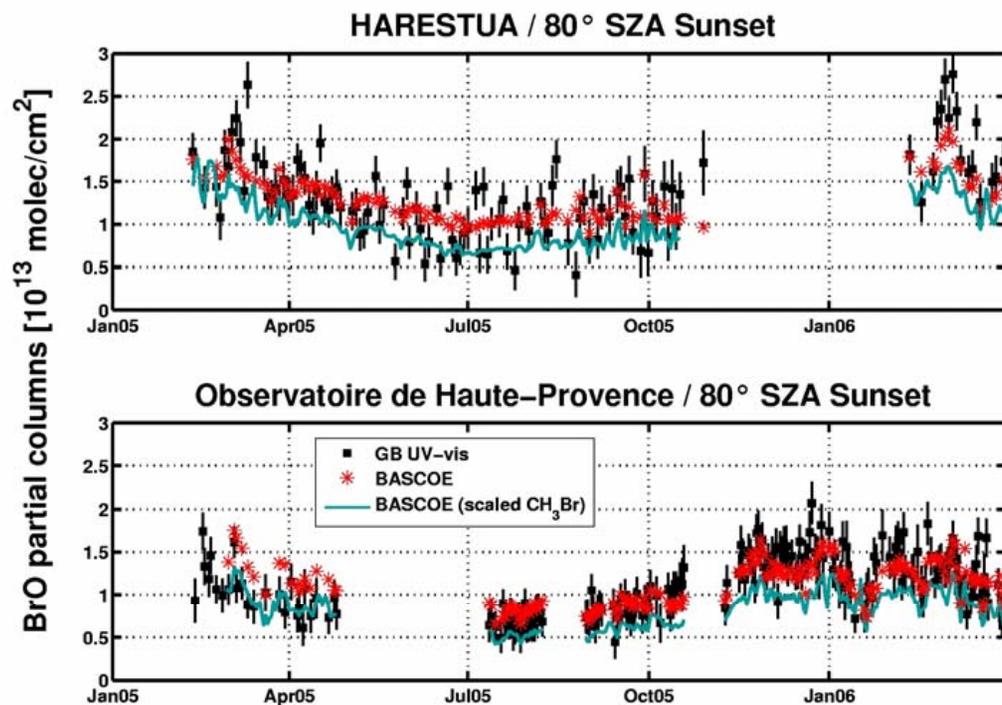


Fig. 9. Comparison between modelled and ground-based stratospheric BrO partial columns (12–20 km altitude range), at Harestua (top) and Observatoire de Haute-Provence (bottom) for the period from January 2005 to March 2006. BASCOE assumes $[Br_y]=23$ pptv for two different scenarios: (red) stratospheric inorganic bromine has contributions from CH_3Br , halons, VLSL and product gases as described in Sect. 2.1., (green) stratospheric Br_y is supplied only by the decomposition of methyl bromide (scaled).

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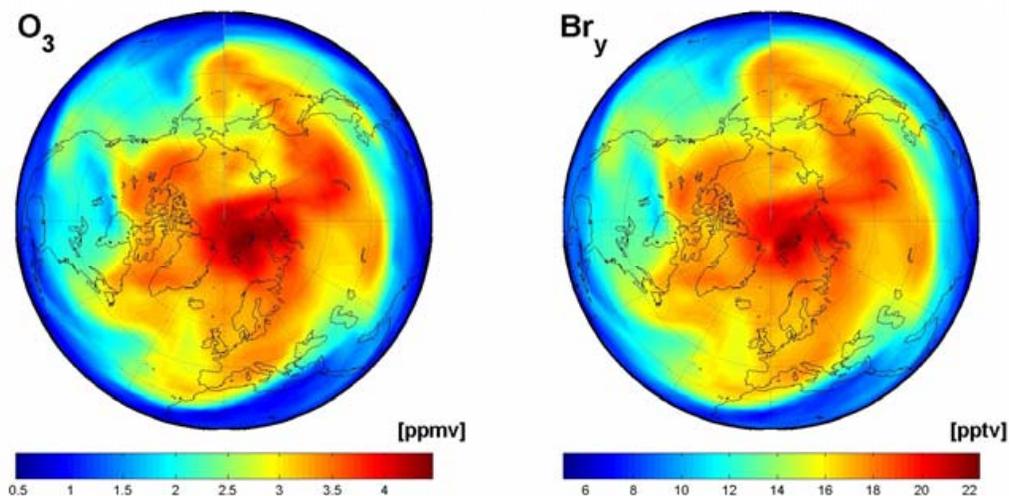


Fig. 10. Example of BASCOE O_3 and Br_y volume mixing ratios in the Northern Hemisphere at a potential temperature of 475°K ($\sim 60\text{ mb}$), for the 1 March 2004 at 00:00 UT.

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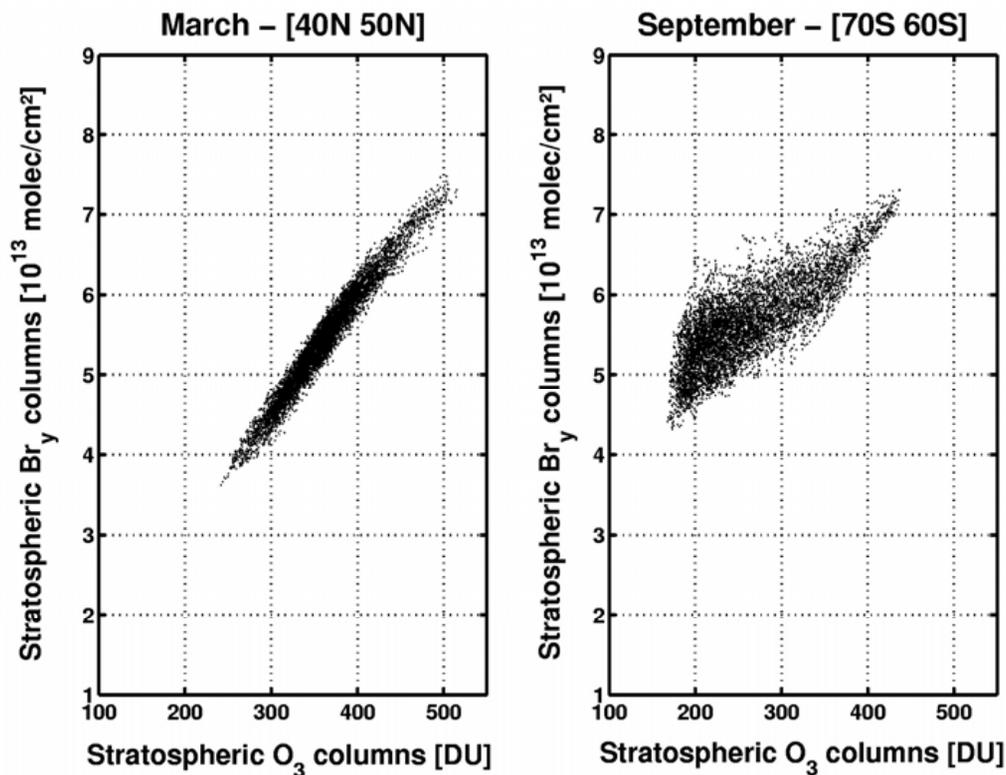


Fig. 11. Correlation between stratospheric Br_y and O₃ vertical columns from BASCOE, for (left) March (40° N–50° N) and (right) September (70° S–60° S).

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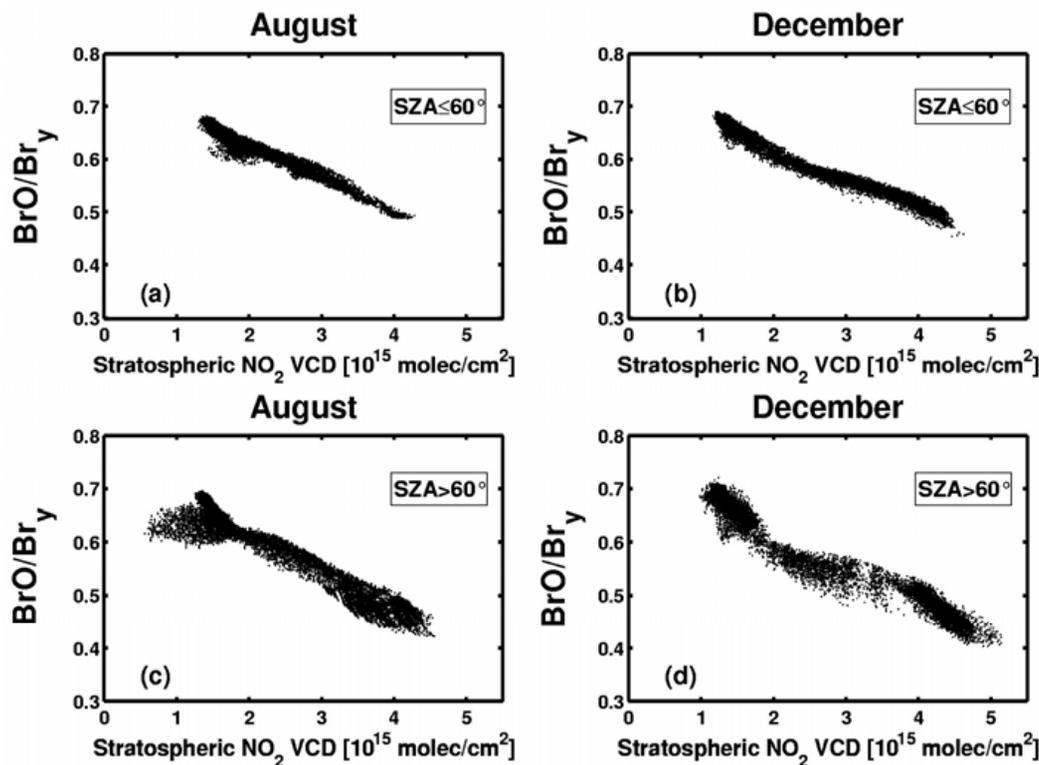


Fig. 12. Ratio between the stratospheric BrO and Br_y columns, as a function of the stratospheric NO₂ columns (morning conditions; upper plots: $SZA \leq 60^\circ$, lower plots: $60^\circ < SZA \leq 80^\circ$), estimated from BASCOE data for August (left plots) and December (right plots).

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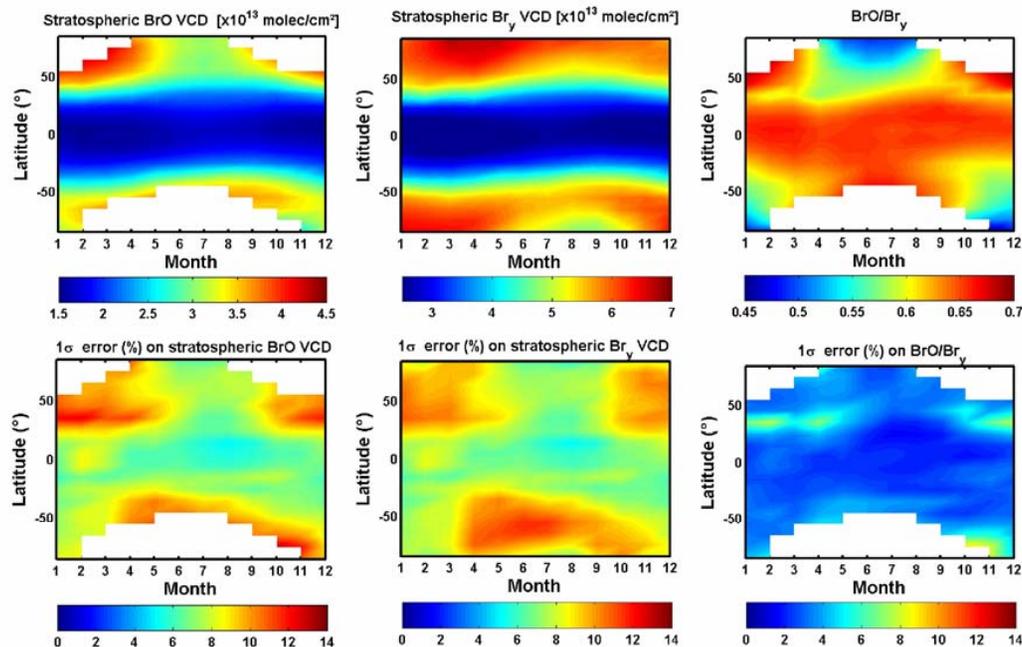


Fig. 13. Results of the BASCOE climatology at GOME overpass. (Upper plots) Monthly zonal averaged of BrO (left) and Br_y (center) stratospheric vertical columns (expressed in $\times 10^{13}$ molec/cm²) and the ratio between BrO and Br_y columns (right). (Lower plots) Standard deviation (expressed in %) on BrO columns climatological values and the respective contributions from the Br_y and BrO/Br_y parameterization. The bromine photochemistry for solar zenith angles larger than 80° is not considered here (white regions).

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