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stratospheric O₃ and
NO₂ abundances**

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Inter-comparison of stratospheric O₃ and NO₂ abundances retrieved from balloon borne direct sun observations and Envisat/SCIAMACHY limb measurements

A. Butz^{1,2}, H. Bösch^{1,*}, C. Camy-Peyret², M. Chipperfield⁵, M. Dorf¹, G. Dufour^{2,**}, K. Grunow⁶, P. Jeseck², S. Kühl¹, S. Payan², I. Pepin², J. Pukite^{1,7}, A. Rozanov³, C. von Savigny³, C. Sioris⁴, T. Wagner¹, F. Weidner¹, and K. Pfeilsticker¹

¹Institut für Umweltphysik, University of Heidelberg, Heidelberg, Germany

²Laboratoire de Physique Moléculaire pour l'Atmosphère et l'Astrophysique (LPMAA), Université Pierre et Marie Curie, Paris, France

³Institute of Environmental Physics and Institute of Remote Sensing, University of Bremen, Bremen, Germany

⁴Harvard-Smithsonian Center for Astrophysics, Cambridge, USA

⁵Institute for Atmospheric Science, School of Earth and Environment, University of Leeds, Leeds, UK

⁶Meteorologisches Institut, Freie Universität Berlin, Berlin, Germany

⁷Institute of Atomic Physics and Spectroscopy, University of Latvia, Riga, Latvia

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* now at: Jet Propulsion Laboratory, California Institute of Technology, Pasadena, USA

** now at: Department of Chemistry, University of Waterloo, Ontario, Canada

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Correspondence to: A. Butz (andre.butz@iup.uni-heidelberg.de)

Abstract

Stratospheric O₃ and NO₂ abundances measured by different remote sensing instruments are inter-compared: (1) Line-of-sight absorptions and vertical profiles inferred from solar spectra in the ultra-violet (UV), visible and infrared (IR) wavelength ranges measured by the LPMA/DOAS (Limb Profile Monitor of the Atmosphere/Differential Optical Absorption Spectroscopy) balloon payload during balloon ascent/descent and solar occultation are examined with respect to internal consistency. (2) The balloon borne stratospheric profiles of O₃ and NO₂ are compared to collocated space-borne skylight limb observations of the Envisat/SCIAMACHY satellite instrument. The trace gas profiles are retrieved from SCIAMACHY spectra using different algorithms developed at the Universities of Bremen and Heidelberg and at the Harvard-Smithsonian Center for Astrophysics. A comparison scheme is used that accounts for the spatial and temporal mismatch as well as differing photochemical conditions between the balloon and satellite borne measurements. It is found that the balloon borne measurements internally agree to within ±10% and ±20% for O₃ and NO₂, respectively, whereas the agreement with the satellite is ±20% for both gases in the 20 km to 30 km altitude range and in general worse below 20 km.

1. Introduction

Stratospheric NO_x (=NO+NO₂) is responsible for up to 70% of the stratospheric O₃ loss (Crutzen, 1970; Portmann et al., 1999). NO_x reactions dominate the catalytic destruction of O₃ between 25 and 40 km altitude via



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Reactions (R1) and (R2), account for more than 90% of NO_x chemistry in the lower stratosphere (Del Negro et al., 1999; Cohen et al., 2000). Thus, NO₂ and O₃ measurements are of primary importance to study the photochemistry of stratospheric O₃. Recent studies by Dufour et al. (2005) indicate that for selected geophysical conditions the agreement between measured and photochemically modeled O₃ and NO_x is better than 10%. Accordingly, high precision measurements are required to constrain or to be compared with photochemical models.

Past observations of these key species involve in-situ as well as optical remote sensing instrumentation operated from ground, aircraft, balloons and satellites, exploiting that O₃ and NO₂ absorb electromagnetic radiation in various wavelength ranges. Historically first the solar occultation technique (e.g. Mauldin et al., 1985; Russell III et al., 1988; Camy-Peyret et al., 1993; Sasano et al., 1993) was applied to the UV/visible and IR spectral ranges and only more recently the satellite-borne UV/visible skylight limb technique became available (e.g. Mount et al., 1984; Rusch et al., 1984; Burrows et al., 1995; von Savigny et al., 2003; Sioris et al., 2003).

The SCanning Imaging Absorption spectroMeter for Atmospheric CHartography (SCIAMACHY) instrument onboard the European Envisat satellite is a UV/visible/near-IR spectrometer designed to measure direct and scattered sunlight in various viewing directions (Burrows et al., 1995; Bovensmann et al., 1999). An exciting new feature of SCIAMACHY is to probe the atmosphere in subsequent and spatially overlapping nadir and limb scanning observations. This will eventually allow to discriminate between the measured total atmospheric column amounts (nadir) and total stratospheric columns obtained from integrated stratospheric profiles to yield tropospheric column amounts of the targeted gases (Sioris et al., 2004).

Here, we present O₃ and NO₂ stratospheric profiles retrieved from SCIAMACHY skylight limb observations using first retrieval exercises developed at the Universities of Bremen (IUP-Bremen) and Heidelberg (IUP-Heidelberg) and at the Harvard-Smithsonian Center for Astrophysics (Harvard). The present study aims at estimating the accuracy of the inferred vertical profiles of stratospheric O₃ and NO₂ by compar-

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ison to the corresponding data retrieved from traditional balloon borne solar occulta-
tion measurements performed by the LPMA/DOAS (Limb Profile Monitor of the Atmo-
sphere/Differential Optical Absorption Spectroscopy) balloon payload (Camy-Peyret,
1995; Ferlemann et al., 1998; Harder et al., 1998; Bösch et al., 2003). For ENVISAT
validation purposes, LPMA/DOAS has been deployed at different launch locations and
in different seasons during the recent past. It allows us to perform simultaneous mea-
surements of targeted gases in various wavelength ranges covering the UV to the IR.

Therefore in a first exercise, the internal consistency of the LPMA/DOAS observa-
tions is checked by comparing slant column amounts of O₃ and NO₂ (taken from the
balloon to the sun) inferred from the visible and IR spectral ranges. Since instrumental
setup and retrieval algorithms are inherently different for the DOAS and LPMA instru-
ment but the line-of-sight is inherently the same, inferred line-of-sight absorptions are
compared and discussed with respect to precision and accuracy of the instruments.
Further, the vertical profiles are analyzed regarding altitude resolution and implications
for satellite validation. In the second part, the balloon borne profiles of O₃ and NO₂
are compared with collocated profiles inferred from SCIAMACHY skylight limb obser-
vations. Spatial and temporal coincidences of the balloon and satellite borne measure-
ments are identified using air mass trajectory calculations based on ECMWF analy-
ses. In addition, for the photochemically short-lived NO₂ radical, the diurnal variation is
modelled on the calculated air mass trajectory in order to consider the different daylight
hour of the satellite and the balloon borne observations. Finally, after accounting for
the spatial and temporal mismatch as well as the differing photochemical conditions,
the balloon and satellite borne measurements are inter-compared and discussed with
respect to inherent errors and possible further improvements of the involved algorithms.

2. Methods

2.1. O₃ and NO₂ profiles inferred from LPMA/DOAS observations

Since details on the instrumental setup and performance of the French/German LPMA/DOAS balloon payload have been reported elsewhere (Camy-Peyret, 1995; Ferlemann et al., 1998), only a short description of the instrumental features important for the present study is given here. The payload is mounted on an azimuth-controlled gondola and comprises a sun-tracker (Hawat et al., 1998) and three optical spectrometers which analyze direct sunlight over virtually the entire wavelength range from the UV to the mid-IR. Sunlight is collected by the automated sun-tracker (beam diameter 10 cm), which points to the center of the solar disk within 30 arcsec. It directs the inner core (beam diameter 5 cm) of the solar beam into a Fourier Transform spectrometer (FT-IR) operated by LPMAA (effective field of view FOV \approx 0.2°) while two small telescopes (diameter 1 cm each, effective field of view FOV \approx 0.53°) mounted into the beam's outer fringe feed the collected sunlight into the two DOAS spectrometers via glass fibre bundles. This optical setup guarantees that the UV/visible (DOAS) and IR (LPMA, FT-IR) spectrometers analyze direct sunlight which traversed almost the same atmospheric air masses, (except for the slightly different effective FOV of both spectrometers). The measurements are performed during balloon ascent/descent and in solar occultation geometry with moderate spectral resolution in the UV/visible (UV: FWHM \approx 0.5 nm, visible: FWHM \approx 1.5 nm) and high spectral resolution in the IR (apodized resolution \approx 0.02 cm⁻¹).

In addition to the spectrometers observing direct sunlight a small versatile UV/visible spectrometer has been operated in limb scattering geometry aboard the same balloon gondola since 2002. The instrumental setup, performance and first results are published in Weidner et al. (2005). The inferred O₃ and NO₂ abundances show very good agreement with the data inferred from the direct sun measurements.

The retrieval of O₃ and NO₂ profiles from LPMA/DOAS measurements is split in two steps. First the trace gas concentrations integrated along the line-of-sight (Slant

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Column Densities (SCDs)) are inferred from individual solar spectra. Then the SCDs are inverted to vertical profiles of the respective trace gases.

2.1.1. DOAS O₃ and NO₂ SCD retrieval:

The UV/visible spectra recorded by the two DOAS spectrometers are analyzed for trace gas absorptions applying the DOAS retrieval algorithm (Platt, 1994; Stutz and Platt, 1996). Each spectrum is evaluated with respect to a solar reference spectrum guaranteeing the removal of solar Fraunhofer lines. The solar reference spectrum usually is the spectrum for which the air mass along the line-of-sight and the residual trace gas absorption are minimal. The residual absorption in the solar reference is determined using Langley's extrapolation to zero air mass.

O₃ SCDs are retrieved from the differential structures in the Chappuis absorption band between 545 nm and 615 nm where the temperature and pressure dependence of the O₃ absorption cross section taken from Anderson et al. (1992) is negligible (Burkholder et al., 1994; Orphal, 2003). Remaining trace gas absorptions are dealt with by simultaneously fitting two NO₂ absorption cross sections recorded at T_≈218 K and T_≈238 K in the laboratory, wavelength aligned to cross sections from Voigt et al. (2002) and orthogonalized with respect to each other. Further, an oxygen dimer (O₄) absorption cross section taken from Hermans (2002) and an H₂O absorption cross section generated from HITRAN 2000 (Rothman et al., 2003) by convolution to our spectral resolution are considered. Rayleigh and Mie scattering are accounted for by including a third order polynomial in the fitting procedure. In addition we allow for a first order polynomial which is subtracted from the measured intensity before any algebraic manipulation to correct for instrumental straylight. In the final evaluation the relative wavelength alignment of the absorption cross sections and the solar reference spectrum is fixed and only the measured spectrum is allowed to shift and stretch.

The line-of-sight absorptions of NO₂ are inferred from the 435 nm to 485 nm wavelength range. Interfering absorption features come from O₄ (Hermans, 2002) and O₃. Two O₃ absorption cross sections recorded in the laboratory at T_≈230 K and T_≈244 K,

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aligned to cross sections from Voigt et al. (2001), are orthogonalized and fitted simultaneously. Broad band spectral features are represented by a fourth order polynomial. Instrumental straylight correction and wavelength alignment are treated as in the case of the O₃ analysis. Additional complications arise from the temperature dependence of the NO₂ absorption cross section while the pressure dependence is negligible at our spectral resolution (Pfeilsticker et al., 1999; Orphal, 2003). The NO₂ analysis is performed using absorption cross sections recorded in our laboratory scaled and aligned to convolved cross sections from Harder et al. (1997) at T=217 K, T=230 K, T=238 K and T=294 K. The resulting four sets of NO₂ SCDs are linearly interpolated to the effective NO₂ concentration weighted temperature along the line-of-sight for each spectrum.

The error bars of the retrieved SCDs are estimated via Gaussian error propagation from the statistical error given by the fitting routine, the error in determining the residual absorber amount in the solar reference spectrum and the errors of the absorption cross sections (1% for O₃, 4% for NO₂). For NO₂ additional errors coming from unaccounted features of the temperature dependence of the absorption cross section (2%) and from the convolution to our spectral resolution (1%) are taken into account. The statistical error comprises the 1- σ fitting error, errors coming from systematic residual absorption features and from shift and stretch of the fitted spectrum. The error of the residual absorber amount in the solar reference spectrum is dominated by the estimated accuracy (5%) of the overhead air mass.

Typical optical densities range between 10⁻¹ and 10⁻³ for both gases. All DOAS data presented in this study originate from spectra in the visible wavelength range. NO₂ SCDs are also retrieved in the 370 nm to 390 nm wavelength range measured by the UV spectrograph. As NO₂ SCDs inferred from the UV and the visible do not differ significantly, only data inferred from the visible spectrograph are shown exhibiting smaller error bars. Evaluating O₃ in the UV is difficult due to the strong temperature dependence and the strong absorption (optical densities \approx 1) below 350 nm which renders the DOAS approach questionable (Frankenberg et al., 2005). For further information on the spectral retrieval see Harder et al. (1998, 2000), Ferlemann et al. (1998, 2000),

and Bösch et al. (2003).

2.1.2. LPMA O₃ and NO₂ SCD retrieval:

The SCD retrieval of O₃, NO₂, NO, HNO₃, N₂O, CH₄, HCl, CO₂ and ClONO₂ is performed simultaneously using a multi fit of 6 to 11 micro-windows. The possibility to retrieve all species depends on the actual filters and beamsplitters used for the FT-IR measurements. The target micro-windows for O₃ and NO₂ are 3040.03 cm⁻¹ to 3040.85 cm⁻¹ and 2914.36 cm⁻¹ to 2915.16 cm⁻¹, respectively. Typically the O₃ absorption lines in the main target window become saturated during deep solar occultation reducing the sensitivity of the retrieval to changes in O₃ abundances along the line-of-sight. Thus, an additional micro-window between 1818.09 cm⁻¹ and 1820.98 cm⁻¹ with non-saturated O₃ absorption features is added if available. Interfering absorbers in the O₃ and NO₂ target windows are H₂O, CO₂, NO, CH₄ and H₂O, O₃, CH₄, respectively. Additional information on ozone SCDs comes from weak absorption in the micro-windows dedicated to NO₂ (2914.36–2915.16 cm⁻¹), and CO₂ (1933.89–1940.00 cm⁻¹). For NO₂, weak absorption in the HCl micro-window (2944.71–2945.11 cm⁻¹) improves the SCD retrieval. Based on absorption line parameters from HITRAN 2004 (Rothman et al., 2005) and a reasonable a priori guess for the trace gas profiles, a forward model calculates synthetic spectra which are fitted to the measured ones by a non-linear Levenberg-Marquardt algorithm. The calculation of the synthetic spectra relies on atmospheric parameters taken from nearby radiosonde launches and climatological and meteorological model data. Fitting parameters include a polynomial of up to third order, a zero order wavelength shift and several parameters to adjust the instrumental line shape (ILS). All auxiliary ILS parameters are determined separately in various test runs and finally set to a fixed value for all spectra during a balloon flight.

The error bars comprise the statistical error of the fitting routine (1-σ), the uncertainty in determining the instrumental line shape, the error coming from the ambient atmospheric parameters and their impact on the spectroscopic parameters and the stated

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error bars of the latter (in total 10% systematic contribution for both gases). For earlier work on LPMA data see [Payan et al. \(1998, 1999\)](#) and [Dufour et al. \(2005\)](#).

2.1.3. LPMA/DOAS O₃ and NO₂ profile retrieval:

Each spectrum yields an O₃ and NO₂ SCD according to the specifications described above. Vertical trace gas profiles are inferred separately from the measurements during balloon ascent and solar occultation. We refer to the respective set of SCDs as the measurement vector \mathbf{y} . In keeping with the standard notation the true trace gas profile is denoted \mathbf{x} ([Rodgers, 2000](#)). The forward model which links the measurements and the vertical profile is straight forward to obtain by raytracing the path of the incoming light from the sun to the detector. Assuming a spherical, layered atmosphere including refraction, the elements K_{ij} of the weighting function matrix \mathbf{K} are given by the ratio of the slant path through layer j and the height of layer j for the observation geometry of spectrum i . The inversion problem can be stated in linear form

$$\mathbf{y} = \mathbf{K}\mathbf{x} + \boldsymbol{\epsilon}, \quad (1)$$

where $\boldsymbol{\epsilon}$ represents the measurement error. A variety of methods exists to invert the weighting function matrix \mathbf{K} and to calculate a retrieved vertical trace gas profile $\hat{\mathbf{x}}$ given the measurements \mathbf{y} . Vertical trace gas profiles from balloon borne measurements shown here are generated using a linear Maximum A Posteriori inversion algorithm described in [Rodgers \(2000\)](#)

$$\hat{\mathbf{x}} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1} (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{y} + \mathbf{S}_a^{-1} \mathbf{x}_a), \quad (2)$$

where the superscript T indicates transposed matrices. A priori, \mathbf{S}_a , as well as measurement covariance matrices, \mathbf{S}_e , are assumed to be diagonal. The a priori profile \mathbf{x}_a is taken from the satellite retrievals described below or from the SLIMCAT chemical transport model ([Chipperfield, 1999](#)). Error propagation is handled by calculating the corresponding error covariance matrix $\hat{\mathbf{S}}$ via

$$\hat{\mathbf{S}} = (\mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K} + \mathbf{S}_a^{-1})^{-1}. \quad (3)$$

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The square roots of the variances \hat{S}_{ij} represent the errors attributed to the retrieved trace gas profile \hat{x} .

The quality of the retrieval can be characterized by the averaging kernel matrix \mathbf{A} which gives the relation between the true trace gas profile \mathbf{x} and the retrieved one \hat{x}

$$\hat{x} = \mathbf{x}_a + \mathbf{A}(\mathbf{x} - \mathbf{x}_a) + \text{error terms.} \quad (4)$$

For the Maximum A Posteriori retrieval \mathbf{A} is given through

$$\mathbf{A} = \hat{\mathbf{S}} \mathbf{K}^T \mathbf{S}_e^{-1} \mathbf{K}. \quad (5)$$

The width of averaging kernels given by the rows of \mathbf{A} is a measure for the altitude resolution of the measurement. As altitude resolution depends on the weighting function matrix \mathbf{K} and measurement error \mathbf{S}_e the altitude resolution of the retrieved trace gas profiles is different for the three sensors presented in this study. When comparing trace gas profiles the differing altitude resolution can be accounted for by degrading the altitude resolution of the high resolution profile. First, the profile inversion of the high resolution measurement is performed on the same altitude grid as the inversion of the low resolution data set, then the resulting trace gas profile \hat{x}_h is smoothed by the averaging kernel matrix of the low resolution measurement \mathbf{A}_l . Accordingly, the smoothed trace gas profile \hat{x}_s is given by

$$\hat{x}_s = \mathbf{x}_a + \mathbf{A}_l(\hat{x}_h - \mathbf{x}_a), \quad (6)$$

where \mathbf{x}_a is the a priori profile simultaneously used for both retrievals (Connor et al., 1994; Hendrick et al., 2004). Whenever smoothed profiles are shown in this study they are generated according to Eq. (6).

2.2. O₃ and NO₂ profiles inferred from Envisat/SCIAMACHY limb observations

The SCIAMACHY instrument, which was put into a sun-synchronous orbit onboard the European Envisat satellite on 28 February 2002, is a UV/visible/near-IR spectrometer

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(220 nm–2380 nm, FWHM: 0.2 nm–1.5 nm) designed to measure either direct sunlight during solar occultation, direct moonlight during lunar occultation or sunlight scattered by the Earth's atmosphere in nadir or limb direction (e.g. [Burrows et al., 1995](#); [Bovensmann et al., 1999](#)). In limb scattering mode SCIAMACHY scans the Earth's atmosphere vertically in steps of 3.3 km from the ground to about 100 km tangent height (vertical field of view at tangent point, $FOV \approx 2.8$ km, horizontal field of view at tangent point, $FOV \approx 110$ km). In addition a horizontal scan is performed at each tangent height covering in total about 960 km at the tangent point. Here, results obtained by the LPMA/DOAS balloon payload since 2003 are used to validate SCIAMACHY limb observations.

O₃ and NO₂ profiles are inferred from SCIAMACHY limb measurements using algorithms developed at the university of Bremen (IUP-Bremen). In the case of NO₂, also data retrieved at the Harvard Smithsonian Center for Astrophysics (Harvard) and a limited data set retrieved by the university of Heidelberg (IUP-Heidelberg) are available. To our knowledge no official ESA products are on hand to be compared with the presented data.

2.2.1. IUP-Bremen O₃ retrieval:

The stratospheric ozone profiles are derived from SCIAMACHY limb scattering observations with the Stratozone code (version 1.62) using the method described in [von Savigny et al. \(2005c\)](#). The altitude range from about 15–40 km can be covered with this technique with a vertical resolution of about 4 km. The retrieval technique exploits the absorption in the Chappuis bands of ozone using only three discrete wavelengths as described in [Flittner et al. \(2000\)](#). A non-linear optimal estimation (OE) scheme drives the radiative transfer model SCIRAYS ([Kaiser and Burrows, 2003](#)), which is run online to calculate weighting functions and to forward-model the limb radiance profiles. The main error sources are incorrect knowledge of the stratospheric aerosol loading, surface albedo, cloud cover, as well as tangent height errors ([von Savigny et al., 2005b](#)). The estimated total error between 15 and 35 km is about 8–14%. The

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SCIAMACHY limb observations are affected by errors in the tangent height information of up to 2.5 km (von Savigny et al., 2005a), particularly before the improvement of the orbit propagator model on-board the Envisat spacecraft in December 2003. Note that tangent height errors of only 500 m lead to errors in the ozone concentrations of up to 10%. For the stratospheric ozone profile retrievals presented here a tangent height offset for each orbit was derived from the limb measurements in the UV-B spectral range prior to the ozone profile retrieval, using the method described in Kaiser et al. (2004). The precision of the tangent height retrieval method is about 200–300 m.

2.2.2. IUP-Bremen NO₂ retrieval:

The forward simulations of the SCIAMACHY limb measurements and the calculations of the appropriate weighting functions are performed employing the SCIATRAN radiative transfer model (Rozanov, 2004; Rozanov et al., 2005). In the spherical mode the SCIATRAN model calculates the limb radiance properly considering the single scattered radiance and using an approximation to account for multiple scattering. Vertical distributions of NO₂ are retrieved from SCIAMACHY limb measurements using the spectral information in the 420–470 nm wavelength interval. To improve the retrieval quality the vertical profiles of O₃ are estimated in combination with NO₂ retrievals using the same spectral information. Limb measurements performed at tangent heights from 12 to 40 km are considered. To reduce the impact of the Fraunhofer structure and incorrect instrument calibration all selected limb scans are divided by the reference limb measurement obtained at a tangent height of about 43 km. To account for broad-band features resulting from unknown scattering properties of the atmosphere as well as instrument calibration issues, a cubic polynomial is subtracted from all spectral ratios. The temperature dependant absorption cross sections of O₃ and NO₂ measured by the SCIAMACHY PFM Satellite Spectrometer are used in the forward model (Bogumil et al., 1999). Pressure and temperature profiles are taken from the corresponding ECMWF data. The SCIAMACHY pointing errors are accounted for by applying appropriate tangent height corrections obtained using the TRUE retrievals (Kaiser et al.,

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2004). The retrieval is performed as described in [Rozanov et al. \(2005\)](#) using a two-step retrieval procedure. During the preprocessing step which is done for each tangent height independently a possible misalignment of the wavelength grids of the limb spectra, of the reference spectrum and of the forward model is accounted for. Additionally, known corrections, namely, undersampling, Ring spectrum, stray light correction and instrument calibration functions, are applied. The main retrieval step is based on the solution of Eq. (1) employing the optimal estimation technique. Different from the description given in Sect. 2.1, the measurement vector \mathbf{y} contains the differences between ratios of simulated and measured differential limb spectra at all tangent heights selected for the retrieval with all corrections from the preprocessing step applied. The state vector \mathbf{x} contains relative differences of trace gas number densities (with respect to a priori values) at all altitude layers for all gases to be retrieved. The final solution is found employing the information operator approach ([Hoogen et al., 1999](#); [Rozanov, 2001](#)) which ensures an additional noise filtering, resulting in more stable profiles.

2.2.3. Harvard NO₂ retrieval:

The Harvard-SAO algorithm is described in [Sioris et al. \(2004\)](#) and references therein. Cloud top height is retrieved from channel 6 and is used to define the lower limit of the retrieval. Tangent height (TH) registration is determined by the multi-wavelength ~305 nm knee method ([Sioris et al., 2003](#)). The calculated orbital median TH offset is applied to all limb scans if greater than the orbital standard deviation of the TH offsets. The NO₂ fitting window consists of 256 pixels in the 434–495 nm range and the simulations are performed at SCIAMACHY channel 3 spectral resolution (0.5 nm), leading to 114 wavelengths in the same fitting window. The analysis uses a classic two-step approach: spectral fitting followed by inversion of SCDs to a number density profile. The reference spectrum is the co-addition of all spectra between the retrieval upper altitude limit (~40 km) and 70 km. Absorption cross sections included in the fitting process are NO₂ and O₃, both at T=203 K, T=223 K, T=243 K from [Bogumil et al. \(2003\)](#), and the collisional oxygen dimer (O₄) from [Greenblatt et al. \(1990\)](#). The

inversion of the SCDs to number density profiles is direct, thus a priori is not required in the retrieval range.

2.2.4. IUP-Heidelberg NO₂ retrieval:

The IUP-Heidelberg uses a two step approach to retrieve NO₂ vertical profiles from SCIAMACHY limb spectra (Kühl, 2005), see also Krecl et al. (2005) and Haley et al. (2004). First, the SCD of the considered absorber is deduced by Differential Optical Absorption Spectroscopy (Platt, 1994). In the second step the SCDs for the different tangent heights are inverted to concentration profiles by an optimal estimation method (Rodgers, 2000). The DOAS retrieval for NO₂ is performed in the 420–450 nm spectral range. The measured and calibrated spectral information from SCIAMACHY is analyzed with respect to a pseudo top-of-the-atmosphere (TOA) reference spectrum taken as average of SCIAMACHY measurements at tangent heights between 40 to 46 km to infer the SCDs. Considered trace gas absorption cross sections are NO₂ at 223 K from Bogumil et al. (2003), O₃ at 241 K from Bogumil et al. (2003), H₂O at 273 K from Rothman et al. (2003), O₄ at 298 K from Greenblatt et al. (1990). In addition we account for the Ring effect (Grainger and Ring, 1962), instrumental straylight and broadband spectral features by considering a calculated Ring spectrum and the inverse of the TOA reference as fitting parameters. The vertical trace gas profiles are obtained using the same approach as in the case of the balloon borne retrieval, see Sect. 2.1. The weighting function matrix **K** is calculated by the fully spherical 3-dimensional Monte Carlo radiative transfer model “Tracy” (von Friedeburg, 2003; Weidner et al., 2005). Using the maximum a posteriori technique the vertical trace gas profile is inferred from Eq. (2), where the measurement vector **y** represents the measured SCDs and the vector \hat{x} the retrieved profile. Our studies show the possibility to retrieve information about the NO₂ concentration in the altitude range from approximately 15 km to 40 km where averaging kernels are larger than 0.7. The accuracy in this altitude range is approximately 15–25%. Tangent height corrections are performed according to the monthly

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averaged SCIAMACHY pointing errors given in [von Savigny et al. \(2005a\)](#).

2.3. Air mass trajectory modelling

Balloon borne measurements exhibit several inherent constraints with respect to the time and location of the balloon launch. Balloon launches are possible at a few sites around the world, only. The launching possibility depends on the local surface weather conditions and the balloon's trajectory is determined by the tropospheric and stratospheric wind fields. Furthermore, the LPMA/DOAS payload is supposed to be launched temporally close to local sunset or sunrise as LPMA/DOAS performs measurements during solar occultation. In practise, these constraints make it difficult to get a direct temporal and spatial coincidence with individual satellite measurements such as SCIAMACHY limb observations. In part, the use of air mass trajectory models can help to overcome the shortcomings in balloon borne satellite validation ([Bacmeister et al., 1999](#); [Lu et al., 2000](#); [Danilin et al., 2002a,b](#)).

For that purpose, an air mass trajectory model is applied within the framework of Envisat/SCIAMACHY validation ([Reimer and Kaupp, 1997](#)). It uses the operational ECMWF analyses and forecasts given every 6 h on a $2.5^\circ \times 2.5^\circ$ latitude/longitude grid. Forward and backward trajectories are calculated on isentropic levels from the surface up to 1600 K with interpolation between the levels. The internal time step is 10 minutes and the diabatic and climatological heating rates are based on Newtonian cooling. The results (trajectory points) are stored for each hour.

Backward and forward trajectories start at the balloon measurement locations. For the remote-sensing payload LPMA/DOAS these starting points are the balloon location and the tangent points for balloon ascent and solar occultation, respectively. For post-flight analysis air mass trajectories are calculated for up to 10 days forward and backward in time, but for balloon flight planning purposes the time range is limited by the available ECMWF forecasts: the latest analysis is from 12:00 UT the day before, forecasts are available every 6 h up to 72 h into the future.

The actual geolocations of Envisat/SCIAMACHY observations are provided

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by the SCIAMACHY Operational Support Team (SOST) through its website (<http://atmos.af.op.dlr.de/projects/scops/>). For each Envisat orbit, overpass time, geolocation and detailed measurement specifications (e.g. swath, measurement duration, ground pixel size) can be downloaded. For the air mass trajectory based matching technique only the area covered by the tangent points of SCIAMACHY limb is considered. This information is used to find satellite measurement points along the individual air mass trajectories, for which the spatial and temporal mismatch is as small as possible. The match criteria are chosen based on the experience of the ozone Match experiment ([von der Gathen et al., 1995](#)). The maximum time mismatch between the satellite observation and the air mass trajectory started at the balloon observation is one hour and the maximum area mismatch is 500 km with respect to the center of the SCIAMACHY limb ground pixel. In case, no Envisat/SCIAMACHY observation satisfies these limits, the distance criterion is weakened to a maximum spatial mismatch of 1000 km.

2.4. Photochemical modelling

On timescales important for this study, i.e. the time between satellite and correlative balloon borne measurement (<1 day), the photochemical variation of O₃ is found negligible. Hence, we focus our modelling on the photochemical variation of NO₂. For the impact of the photochemistry of NO₂ on solar occultation measurements and validation studies see also [Kerr et al. \(1977\)](#), [Roscoe and Pyle \(1987\)](#) and [Bracher et al. \(2005\)](#). We use output from a simulation of the SLIMCAT 3-D off-line chemical transport model (CTM) ([Chipperfield, 1999](#)) to initialize a 1-D chemistry model of the stratosphere. SLIMCAT output of run 323 is saved at 00:00 UT every 2 days interpolated to the launching sites of the balloon flights. Both photochemical models include a comprehensive set of the relevant gas-phase and heterogenous reactions as given by the JPL-2002 report on Chemical Kinetics and Photochemical Data ([Sander et al., 2003](#)). The 1-D chemistry model is an updated version of the model used in [Bösch et al. \(2003\)](#). Stratospheric chemistry is modelled on 19 potential temperature (Θ) levels

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between $\Theta=336\text{ K}$ ($\approx 11\text{ km}$) and $\Theta=1520\text{ K}$ ($\approx 42\text{ km}$). Aerosol loadings are taken from [Deshler et al. \(2003\)](#) as recommended by [Dufour et al. \(2005\)](#). Photolysis rates are interpolated with respect to pressure, temperature, overhead ozone and solar zenith angle (SZA) from a lookup table where the actinic fluxes are calculated as recommended by [Lary and Pyle \(1991\)](#) and validated for J_{NO_2} by [Bösch et al. \(2001\)](#).

If available the 1-D model is initialized at 00:00 UT with SLIMCAT output of the same day at the balloon launch site. If output is not available on the day of the balloon flight, we decide whether to take output from the day before or after the flight by comparing the measured O₃, NO₂ and if available CH₄ and N₂O profiles to the modelled ones and choosing the output which produces better agreement with the measurement. While the model is run with fixed pressures and temperatures for each Θ level taken from the meteorological support data of the balloon flight, the SZA timeline is taken from the air mass trajectory calculations to guarantee that the photochemical evolution of the modelled air mass corresponds to the true evolution between initialization of the model, satellite measurement and balloon borne observation. For simplicity a single representative SZA timeline is chosen for all Θ levels. In addition O₃, NO₂, NO and N₂O₅ are scaled at initialization in a way that the model can reproduce the balloon borne O₃ and NO₂ profiles in the evening ([Bracher et al., 2005](#)).

In general, each spectrum measured by SCIAMACHY as well as by the balloon borne instruments is a composite of several photochemical conditions since the SZA varies along the line-of-sight through the atmosphere. All satellite measurements presented here are conducted in the morning far from sunset or sunrise where the photochemical variation of NO₂ is weak. Hence, we assume a fixed SZA for the SCIAMACHY observations. However, the time-lag to the validation measurements is on the order of several hours and balloon borne profiles are inferred from measurements close to sunset when the photochemical variation of NO₂ is strongest. Introducing the photochemical weighting factors κ_{kj} , balloon observations are scaled to the photochemical conditions of the satellite measurements. Given the model data along the SZA timeline

of a given air mass trajectory, κ_{kj} is defined by

$$\kappa_{kj} = \frac{a_{kj}}{b_j}, \quad (7)$$

where b_j is the modelled NO_2 concentration at altitude j and the SZA of the SCIAMACHY measurement and a_{kj} the modelled NO_2 concentration at altitude j and SZA k . When tracing the light through the atmosphere from the sun to the balloon borne detector, each point on the line-of-sight can be identified through its altitude and its local SZA. Assuming a layered atmosphere the slant path through layer j at local SZA k is multiplied by κ_{kj} . Hence, we obtain a photochemically corrected weighting function matrix $\tilde{\mathbf{K}}$. Replacing \mathbf{K} by $\tilde{\mathbf{K}}$ in Eq. (1) and solving the inversion problem as described in Sect. 2 yields the balloon borne trace gas profile scaled to the photochemical conditions of the satellite measurement.

Following a similar approach as in Bracher et al. (2005) the modelling error is estimated by sensitivity studies. For the flight from Kiruna on 24 March 2004, several model runs are performed along a representative air mass trajectory varying model parameters important for the photochemical variation of NO_2 . These parameters include the O_3 profile (-33% , $+50\%$), overhead ozone ($\pm 74\%$), the temperature for each Θ -level ($\pm 7\text{ K}$), the rate constants of Reactions (R1) and (R2) ($\pm 20\%$), the aerosol surface area ($\pm 40\%$) and the γ -coefficient for N_2O_5 uptake on liquid aerosol (-50% , $+100\%$). For each run two photochemically scaled trace gas profiles are generated corresponding to a backward and a forward satellite match in the morning before and after the balloon flight, respectively. The root mean square deviation of the vertical profiles from the standard run yields the estimate of the modelling error. The modelling error increases from less than 10% at balloon float altitude to 20% at 20 km altitude. Between 20 km and balloon float altitude the error is governed by the factors which influence the photolysis of NO_2 via Eq. (R2). The modelling error for the backward match grows to 30% at 15 km and up to 50% below 15 km, mainly due to the sensitivity of NO_x chemistry to the O_3 profile through Eq. (R1). The modelling error of the forward match calculations

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remains constant at 20% down to 15 km. Below 15 km it rises up to 35% caused by sensitivity to the O_3 profile. Whenever photochemically corrected trace gas profiles are shown the modelling error is added applying Gaussian error propagation.

3. Internal LPMA/DOAS comparison

3.1. Comparison of LPMA/DOAS O_3 and NO_2 Slant Column Densities

Table 1 summarizes the geophysical conditions of the balloon flights chosen for the comparison of the LPMA/DOAS O_3 and NO_2 SCDs. We present data of six joint LPMA/DOAS balloon flights where the corresponding geophysical conditions range from high-latitude winter and summer to mid-latitude fall comprising measurements inside as well as outside the northern polar vortex. Due to bad data quality or instrumental malfunction and the considerable effort necessary for data re-analysis we restrict the comparison study to six out of the 13 flights conducted to date.

The quality of the LPMA/DOAS comparison is illustrated in Fig. 1 showing O_3 and NO_2 SCDs as well as the corresponding errors and relative deviations inferred from observations at Kiruna on 23 March 2003. The particular flight is chosen as an example since all data sets are available allowing for a consistent comparison. In addition, relative deviations between LPMA and DOAS SCDs are presented as histograms in Figs. 2 and 3 for the six considered balloon flights. Relative deviations are calculated by linear interpolation of the DOAS data to the measurement instances of LPMA.

The general agreement of LPMA and DOAS O_3 SCDs is on the order of 10%. Maximum deviations of up to 30% are observed when slant columns are very low and governed by the O_3 abundances above balloon float altitude indicating a problem when applying Langley's method to retrieve the residual absorber amount in the DOAS solar reference spectrum. For some of the earlier balloon flights the relative deviations for solar occultation measurements are as large as 20%. During solar occultation the measured O_3 IR-absorption lines in the LPMA standard retrieval window become sat-

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urated. In order to increase retrieval sensitivity we include a micro-window exhibiting weak, unsaturated O₃ absorption. Unfortunately the latter micro-window is not available for some of the flights due to usage of spectral filters which cut the respective spectral region. The statistical analysis of all 1032 data points yields that in average DOAS O₃ SCDs are larger by 6.1% than the LPMA data. The standard deviation of the relative deviation between LPMA and DOAS amounts to 8.0%, see Fig. 2.

The comparison of NO₂ SCDs is more difficult to assess. The NO₂ measurements during balloon ascent of the earlier flights from Leon in 1996 and Kiruna in 1997 are hard to compare as LPMA data are very noisy. Data retrieved from balloon ascent of the more recent flights at Kiruna in 2001 and in 2003 are less noisy although the NO₂ slant column abundances show similar values. This suggests that the preflight optical alignment of the FT-IR improved during the series of presented balloon flights. For the flights at Kiruna and at Gap in 1999, NO₂ slant columns are close to the detection limit of the LPMA retrieval causing large error bars. The agreement on NO₂ SCDs inferred from solar occultation is on the order of 20%. The maximum deviation up to 50% occurs for the flight from Kiruna in 1999 where NO₂ abundances are very low. In this case, stratospheric temperatures were well below 217 K where NO₂ absorption cross sections from Harder et al. (1997) are not available. Thus, the DOAS NO₂ data have to be extrapolated when adjusting the retrieved SCDs to the effective mean temperatures along the line-of-sight. The extrapolation might not reproduce the true atmospheric state. The statistical analysis of all 1016 data points reveals a mean deviation of +3.7% between DOAS and LPMA with a large standard deviation of 24.5%, see Fig. 3 lower panel. As some of the LPMA data are noisy or close to the detection limit we performed another statistical analysis excluding all data where the corresponding SCDs show errors larger than 25%. The resulting set of 753 data points exhibits a mean relative deviation between DOAS and LPMA of +6.6% and 14.0% standard deviation, see Fig. 3 upper panel. Overall the combined error bars are a reasonable estimate for the LPMA/DOAS agreement on NO₂ measurements.

3.2. Comparison of LPMA/DOAS NO₂ profiles

Section 3.1 discussed the level of agreement between the O₃ and NO₂ SCDs inferred from LPMA and DOAS measurements. In this section NO₂ data again taken from balloon ascent and sunset of the flight at Kiruna on 23 March 2003, are chosen to illustrate the characteristics of the profile inversion technique. SCDs are inverted to vertical trace gas profiles via Eq. (2). The profile retrieval is characterized by the corresponding averaging kernel matrix given by Eq. (4). The NO₂ profiles presented in Fig. 4a and b for balloon ascent and sunset, respectively, are retrieved using a diagonal a priori covariance matrix with variances corresponding to 30% error of the a priori profile. LPMA data are inverted on a 2-km altitude grid. DOAS data are generated on the same altitude grid as the satellite retrievals which corresponds to 1 km grid spacing between 10 and 45 km altitude and coarser below and above. In addition DOAS profiles are shown which are smoothed applying Eq. (6) to match the coarser altitude resolution of LPMA. The corresponding averaging kernels are shown in Fig. 4c and d for balloon ascent and sunset, respectively. For comparison, the averaging kernels of the NO₂ SCIAMACHY limb retrieval (IUP-Bremen) are also plotted.

The retrieved NO₂ profiles reproduce the general behavior of the underlying SCDs shown in Fig. 1c and d. LPMA data underestimate the DOAS data during balloon ascent. Below 22 km the underestimation is substantial as already indicated by the corresponding SCDs, above 22 km the agreement is better and mostly on the order of the combined error bars. For sunset, LPMA overestimates the DOAS data between 31 km and 25 km. Below 25 km LPMA and DOAS agree well.

The averaging kernels indicate that it is possible to retrieve a vertical profile in the range between 11 km and 29 km exhibiting 5 km to 1 km altitude resolution from DOAS measurements during balloon ascent. LPMA measurements during balloon ascent allow for profile retrieval between 17 km and 29 km altitude with 5 km to 2 km altitude resolution. Observations in solar occultation geometry during sunset yield altitude information in the 32 km to 12 km and 32 km to 18 km range for DOAS and LPMA, re-

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spectively. The corresponding altitude resolution decreases from 1 km at the top to about 4 km at 12 km altitude for DOAS, and from 2 km at the top to about 5 km at 18 km altitude for LPMA. The decrease in altitude resolution is mostly due to the apparent size of the solar disk getting larger along the line-of-sight when penetrating deeper into the atmosphere. Comparing the averaging kernels of the two data sets it becomes evident that smoothing is most significant below 20 km altitude where LPMA measurements yield only little or no information on the NO₂ vertical profile for the selected balloon flight. The secondary (and real) maximum at about 15 km altitude seen in the DOAS unsmoothed data is not present in the LPMA retrieval. Hence, according to Eq. (6) the smoothed DOAS profile is drawn towards the a priori. SCIAMACHY averaging kernels are smaller than 0.4 and significantly broader than the grid spacing indicating that on the selected retrieval grid the retrieved profile points will not be independent. Correspondingly the altitude resolution amounts to 3 km to 5 km in the retrieved altitude range.

3.3. Conclusions on the LPMA/DOAS comparison and implications for satellite validation

As presented in Sects. 3.1 and 3.2, O₃ and NO₂ abundances inferred from LPMA/DOAS measurements in the IR and visible spectral ranges agree roughly within ±10% for O₃ and ±20% for NO₂. Observed discrepancies can be due to instrumental and retrieval related shortcomings of either technique. The LPMA Fourier Transform spectrometer inherently exhibits smaller signal-to-noise-ratio (SNR_≈10²) than the grating spectrometers operated by DOAS (SNR_≈10⁴) causing significantly smaller detection limit and higher precision of the DOAS measurements. In the visible spectral range where the Sun's intensity peaks the typical integration time for individual spectra is less than 1 s, whereas in the IR it takes about 50 s to record a single interferogram. Hence, the DOAS instruments sample the atmosphere with a much higher rate than the LPMA FT-IR. Taking into account the respective errors of the considered trace gases it is evident that altitude resolution is significantly better for the DOAS observations.

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When the apparent size of the solar disk becomes large during solar occultation the smaller field of view of the FT-IR ($FOV \approx 0.2^\circ$) partly compensates the high integration times and large errors compared to the DOAS measurements ($FOV \approx 0.53^\circ$). On the other hand the small field of view of the FT-IR renders the intensity of the measured interferograms very sensitive to small pointing errors of the sun-tracker as the solar irradiance is not uniform across the solar disk (Bösch et al., 2003). According to Fig. 2 and 3 LPMA O₃ and NO₂ SCDs are low biased by 6.1% and 6.6% with respect to the corresponding DOAS SCDs. For both trace gases the typical errors of individual measurements are larger than the bias. Errors of the absorption cross sections with respect to magnitude or pressure and temperature dependencies would directly cause a bias. Especially the DOAS retrieval of NO₂ might exhibit shortcomings when accounting for the temperature dependence of the NO₂ absorption cross section, since the interpolation is performed linearly using four reference temperatures, only. In the IR spectral range, errors in the spectroscopic parameters and the ambient temperature and pressure data propagate into the shape of the absorption lines causing systematic errors of the retrieved trace gas SCDs. Correlations between the various fitting parameters considered for the spectral retrieval are minimized but, still, might cause small unaccounted systematic errors in both data sets. Summing up the arguments, the O₃ and NO₂ stratospheric profiles inferred from DOAS measurements appear to be better suited for validation purposes than the LPMA data particularly as the former exhibit smaller error bars and better altitude resolution. Nonetheless, stratospheric FT-IR measurements are indispensable since the majority of trace gases important for a comprehensive understanding of stratospheric photochemistry do not absorb in the UV/visible.

4. Validation of SCIAMACHY O₃ and NO₂ profiles

4.1. Observations

The validation study reports on four LPMA/DOAS balloon flights, three from ES-RANGE, Kiruna, Sweden, and one from Aire sur l'Adour in southern France conducted since 2003. The analysis of a fifth balloon flight at Teresina, Brazil, in June 2005 is currently in preparation. For each balloon flight a satellite coincident measurement is identified in the morning before and after the balloon flight using the trajectory matching technique described in Sect. 2.3. In the following we refer to these coincidences as backward and forward coincidences. If trace gas profiles inferred from balloon ascent and solar occultation are available the satellite coincidences are identified separately. For each balloon flight Table 2 provides information on the measurement site, the geophysical condition, the SZA range covered by the balloon borne observations, the available data sets and some details on the selected SCIAMACHY limb scans. Besides orbit number, time and date of the satellite measurement, the matching altitude, the distance with respect to the air masses probed by the balloon borne instruments and the time delay between satellite overpass and balloon flight are given.

In general we succeed in identifying satellite coincidences in the altitude range from about 20 km to 30 km. In some cases the characteristics of the satellite coincidences are not perfect for validation purposes, e.g. for the flight from Kiruna in 2004 where the coincident altitude range is below 20 km and for the flight from Kiruna on 4 March 2003, where the forward coincidence is restricted to a small altitude range. Nevertheless, remembering that coincidence criteria are somewhat artificial the balloon and satellite borne trace gas profiles are compared in the whole retrieved altitude range. For the flight from Aire sur l'Adour on 9 October 2003, the match criterion is weakened to a maximum spatial mismatch of 1000 km, as no coincidence is found applying the 500 km criterion. Furthermore assuming the weakened match criterion, the satellite overpass in the morning of 9 October coincides directly with the balloon measurement locations. Hence, the backward model calculations for this flight are not performed along the SZA

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timeline given by air mass trajectory calculations, but are carried out along the SZA timeline corresponding to air masses stationary at Aire sur l'Adour.

The retrieved vertical trace gas profiles for O_3 and NO_2 are shown in Fig. 5 and Fig. 6, respectively. Satellite observations in the morning before and after the balloon flight are compared to balloon borne trace gas profiles. We present smoothed and in the case of NO_2 also photochemically corrected balloon borne profiles. For reference vertical profiles without photochemical correction at full altitude resolution are plotted. In the case of O_3 , in-situ sonde data are included if available. The O_3 sonde data are taken from an electrochemical cell either deployed onboard the gondola or launched from the ground station shortly after balloon launch. The relative deviations between the satellite and balloon borne measurements are assessed in Fig. 7 and Fig. 8 for O_3 and NO_2 , respectively.

LPMA data are available for the two flights at Kiruna in 2003, only. DOAS data are lacking for 4 March 2003, due to instrumental malfunction. According to the findings of Sect. 3.2, DOAS O_3 and NO_2 profiles are retrieved on the same altitude grid as the satellite profiles. As the altitude resolution of the DOAS retrieval is significantly better, DOAS profiles are smoothed to match the satellite's coarser altitude resolution except for NO_2 at balloon ascent below 20 km where no smoothing is applied. LPMA trace gas profiles are generated on a 2-km altitude grid applying no smoothing as the satellite's altitude resolution does not differ significantly. Satellite borne O_3 profiles are inferred using the IUP-Bremen retrieval algorithm. SCIAMACHY NO_2 profiles are generated by the IUP-Bremen and Harvard retrieval. In two cases also NO_2 profiles inferred by the IUP-Heidelberg are available. Relative deviations between balloon and satellite borne observations are calculated for the trace gas profiles retrieved at the IUP-Bremen and the balloon borne DOAS measurements, except for the observations at Kiruna on 4 March 2003, where LPMA data are used.

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4.2. Discussion

4.2.1. O₃ validation study:

The internal consistency of the validation data set has been discussed in Sect. 3.1 where a bias of +6.1% between DOAS and LPMA O₃ measurements is detected. The corresponding standard deviation of the deviations between the two datasets is 8.0%. Accordingly, the inferred vertical profiles mostly agree to within the combined error bars of both data sets. In all cases the agreement between in-situ sonde and remote sensing balloon borne O₃ data is good. Sometimes even highly filamented structures, e.g. Fig. 5e, can be observed simultaneously in the high resolution balloon borne and the in-situ observations.

Figure 5 and Fig. 7 show the O₃ satellite validation study. In most cases SCIAMACHY limb O₃ profiles agree to within $\pm 20\%$ with the validation data set in the 20 km to 30 km altitude range. The agreement ranges from close to perfect as for the observations at Kiruna on 4 March 2003, Fig. 5a, to fair as in the case of Fig. 5d, Aire sur l'Adour on 9 October 2003, where the agreement in the considered altitude range is on the order of 30%, only, and the profile shape is rather different for satellite and balloon borne data. The relative deviations show a systematic underestimation of the balloon borne by the satellite borne profiles at 24 km to 28 km altitude. The underestimation decreases and sometimes changes to overestimation when going up to 31 km and down to 20 km. This finding is similar to conclusions of Brinksma et al. (2005), where a zigzag shape of the deviations between the validation data set (lidar, SAGE II) and the IUP-Bremen O₃ retrieval is observed, indicating that O₃ concentrations at 27 km inferred from SCIAMACHY limb are too low. Albeit different corrections for tangent height errors are already included in the SCIAMACHY retrievals there might be a remaining small tangent height error causing the observed deviations.

Below 20 km SCIAMACHY O₃ profiles underestimate the balloon borne data in most cases and cannot reproduce the frequently highly filamented O₃ profiles observed especially at high latitudes during winter, e.g. Fig. 5e and f. The balloon flights from Kiruna

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on 23 March 2003, and on 24 March 2004, have been conducted close to the polar vortex edge where the gradients in O₃ concentration are large on small spatial scales. When identifying coincident balloon and satellite measurements, the air mass trajectory calculations allow for a sizeable mismatch in space and time. Spatial mismatch of up to 500 km is possible and, hence, the influence of horizontally inhomogeneous air masses can be important close to the polar vortex edge. Further, SCIAMACHY measurements represent an average over a 960 km wide horizontal area, whereas the LPMA/DOAS measurements average the horizontal trace gas distribution along the line-of-sight from the balloon to the sun. Deviations as in Fig. 5e and f might originate from the difference in horizontal averaging. Good agreement is found for the backward coincidences in Fig. 5b and c. However, the forward coincident measurements in Fig. 5b, c, e and f show large deviations between satellite and balloon borne data. Since the uncertainty of the air mass trajectory calculations increases with increasing time-lag to initialization of the trajectory model, the influence of horizontally inhomogeneous air masses might be enhanced for the forward calculations where the time-lag is about ten hours larger than in the backward case. Unfortunately, there are no DOAS or LPMA data available for the observations below 19 km at the mid-latitudinal station Aire sur l'Adour, but SCIAMACHY limb profiles and the corresponding profiles inferred from in-situ O₃ sondes agree well, see Fig. 5d.

4.2.2. NO₂ validation study:

The internal agreement of the balloon borne NO₂ measurements has been assessed in Sects. 3.1 and 3.2. We observe a bias of +6.6% of the DOAS with respect to the LPMA observations, the standard deviation of the relative differences is 14.0%. Accordingly, the LPMA/DOAS observations in Fig. 6b and c mostly lie within the combined error bars. Nonetheless, for the flight from Kiruna on 23 March 2003, LPMA underestimates the DOAS profile for balloon ascent and overestimates it for sunset as already seen in Fig. 4. Remarkable agreement can be observed for DOAS profiles retrieved from measurements during balloon ascent and sunset which are scaled to the photochemical

regime of the same SCIAMACHY observation, proofing the validity of our photochemical as well as meteorological approach.

The comparison between NO₂ profiles inferred from the SCIAMACHY limb measurements and the validation data set is presented in Fig. 6 and Fig. 8. The internal agreement of the satellite retrievals is variable. NO₂ profiles inferred by IUP-Bremen and IUP-Heidelberg show good agreement. Deviations are observed for 9 October 2003, Fig. 6d, at 30 km altitude and for 24 March 2004, Fig. 6e, below 20 km. In some cases, e.g. Fig. 6d, the Harvard NO₂ retrieval yields smaller trace gas concentrations than the IUP-Bremen algorithm around 25 to 27 km altitude. In one case, the backward coincidence in Fig. 6e and f, where SCIAMACHY was operating in a non-nominal measurement mode and pointing could not be verified, the Harvard profiles are offset by +2 km to +3 km with respect to the other satellite data. For the forward coincidence at Kiruna in 2004, agreement is good. However, a detailed comparison of the different retrieval exercises is beyond the scope of this study. In the following we refer our discussion to the IUP-Bremen records as the latter agree best with the balloon borne data. Implications for the Harvard and IUP-Heidelberg retrievals can be inferred easily.

In the 20 km to 30 km altitude range the agreement between the balloon borne NO₂ profiles and the satellite observations is on the order of 20%. No clear trend can be observed except for the backward observations at Kiruna on 4 March 2003, Fig. 6a, where an altitude offset of +2 km to +3 km of the SCIAMACHY with respect to the LPMA data is clearly visible. The same might be true for the forward coincidences at Kiruna in 2004, Fig. 6e and f, but the offset is not as clear as for the former comparison.

Below 20 km the level of agreement is variable. Similar to the O₃ comparison, the backward coincidences in Fig. 6b, c, e and f reveal moderate deviations whereas the corresponding forward coincidences exhibit larger differences between the satellite and the balloon borne measurements. In the case of NO₂ an additional source of uncertainty at low altitudes is the increasing error of the photochemical model calculations which is discussed in Sect. 2.4. Further, as can be seen in Fig. 6e, the SCIAMACHY limb profiles retrieved by the IUP-Bremen, IUP-Heidelberg and Harvard algo-

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rithms exhibit sizeable discrepancies below 20 km. This indicates that for low altitudes the SCIAMACHY retrieval might depend on the actual parameters, e.g. a priori information, used. The latter finding is supported by the characteristics of the corresponding averaging kernels, Fig. 4, which grow wider below 20 km altitude.

5. Conclusions

Stratospheric O₃ and NO₂ abundances inferred from different sensors are inter-compared.

In the first part of this study line-of-sight absorptions and vertical profiles retrieved from the UV/vis DOAS spectrometer and the LPMA FT-IR both performing balloon borne direct sun measurements during balloon ascent and solar occultation are compared. The general agreement is $\pm 10\%$ and $\pm 20\%$ for O₃ and NO₂, respectively. The observations in the visible wavelength range exhibit higher precision and better altitude resolution than the FT-IR measurements due to lower instrumental noise and higher sampling frequency. In solar occultation, the smaller field of view of the FT-IR partly compensates the deficiency in altitude resolution. A small bias between the data inferred from the visible and from the IR observed for both gases could be explained by errors of the spectroscopic parameters, i.e. absorption cross sections and their pressure and temperature dependencies, or remaining correlations between the retrieval parameters. However, the bias lies within the errors of individual measurements.

The second part of this study addresses the validation of O₃ and NO₂ stratospheric profiles inferred from SCIAMACHY skylight limb observations based on the balloon borne data set presented in the first part. An air mass trajectory model is used to identify coincident balloon and satellite borne measurements. The balloon borne trace gas profiles are treated appropriately to match the altitude resolution of the satellite sensor and in the case of NO₂ by means of a 1-D stratospheric chemistry model to match the photochemical conditions of the satellite measurements. Typical deviations between SCIAMACHY observations and balloon borne data amount for both consid-

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ered gases to 20% in the 20 km to 30 km altitude range, somewhat depending on the retrieval algorithm. In the case of O₃ our observations support findings of a previous study (Brinkma et al., 2005) indicating that the IUP-Bremen O₃ retrievals are systematically low between 24 km and 28 km altitude. Below 20 km the agreement worsens for both gases due to the lower sensitivity of the satellite retrieval, uncompensated horizontal variations of the trace gases and in the case of NO₂, modelling uncertainty. The internal agreement of the satellite NO₂ retrieval exercises developed by IUP-Bremen, Harvard and IUP-Heidelberg, is promising. Finally, the present study provides a data set which can contribute to the improvement and the validation of future official ESA algorithms currently under development.

The presented balloon borne profiles will be made available to the public via our web-site <http://www.iup.uni-heidelberg.de/institut/forschung/groups/atmosphere/stratosphere/>.

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Table 1. Compendium of joint LPMA and DOAS balloon flights chosen for comparison. BA, BD, SS and SR indicate balloon ascent, balloon descent, sunset and sunrise, respectively.

Balloon flight date, time/UT	Location	Geophysical condition	Observation geometry
23 Nov. 1996 14:56–16:54	Leon 42.6° N, 5.7° W	mid-lat. fall SZA: 74.0°–92.7°	BA, SS
14 Feb. 1997 12:29–14:53	Kiruna 67.9° N, 21.1° E	high-lat. winter SZA: 82.7°–94.4°	BA, SS
10 Feb. 1999 11:40–14:52	Kiruna 67.9° N, 21.1° E	high-lat. winter SZA: 82.8°–94.6°	BA, SS
25 June 1999 04:32–07:46	Gap 44.2° N, 4.6° E	mid-lat. summer SZA: 86.7°–54.7°	BD
21/22 Aug. 2001 16:04–20:05 01:15–02:30	Kiruna 67.9° N, 21.1° E	high-lat. summer SZA: 75.6°–95.8° SZA: 95.4°–90.0°	BA, SS SR
23 Mar. 2003 14:47–17:28	Kiruna 67.9° N, 21.1° E	high-lat. spring SZA: 77.9°–94.7°	BA, SS

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Table 2. Compendium of joint LPMA and DOAS O₃ and NO₂ profile measurements and Envisat/SCIAMACHY overpasses. BA and SO indicate balloon ascent and solar occultation measurements, respectively.

Balloon flight date, time/UT	Location	Geophysical condition	Available datasets	Satellite coincidence orbit, date, time/UT	Altitude range/km	Time delay/h	Spatial distance/km
04 Mar. 2003 13:20–16:17	Kiruna 67.9° N, 21.1° E	high lat. spring SZA: 71.1°–94.1°	SO: LPMA	5273, 4 Mar., 11:05	20–30	–5.1	369–496
				5285, 5 Mar., 7:17	23–24	+15.3	498–499
23 Mar. 2003 14.47–17:28	Kiruna 67.9° N, 21.1° E	high lat. spring SZA: 78.9°–94.7°	BA: LPMA, DOAS SO: LPMA, DOAS	5545, 23 Mar., 11:07	18–28	–5.2	268–496
				5558, 24 Mar., 9:01	19–29	+17.4	10–495
				5545, 23 Mar., 11:07	20–30	–6.2	63–458
				5558, 24 Mar., 9:01	17–30	+16.0	256–453
9 Oct. 2003 15:39–17:09	Aire sur l'Adour 43.7° N, 0.3° W	mid-lat. fall SZA: 72.0°–87.8°	BA: DOAS	8407, 9 Oct., 9:51	17–31	–6.5	738–988
				8421, 10 Oct., 9:20	25–33	+17.2	547–977
24 Mar. 2004 14:04–17:31	Kiruna 67.9° N, 21.1° E	high lat. spring SZA: 74.5°–95.3°	BA: DOAS SO: DOAS	10798, 24 Mar., 10:35	12–33	–5.4	371–499
				10812, 25 Mar., 10:04	6–16	+19.9	32–485
				10798, 24 Mar., 10:35	10–33	–6.9	191–436
				10812, 25 Mar., 10:04	10–20	+16.7	301–475
17 June 2005 18:32–21:13	Teresina 5.1° S, 42.9° W	tropical winter SZA: 60.6°–95.8°	BA: in preparation SO: in preparation	17240, 17 June, 11:53	25–30	–8.1	382–491
				17255, 18 June, 13:02	5–33	+18.4	6–490
				17240, 17 June, 11:53	23–32	–9.1	519–971
				17255, 18 June, 13:02	8–33	+16.2	12–496

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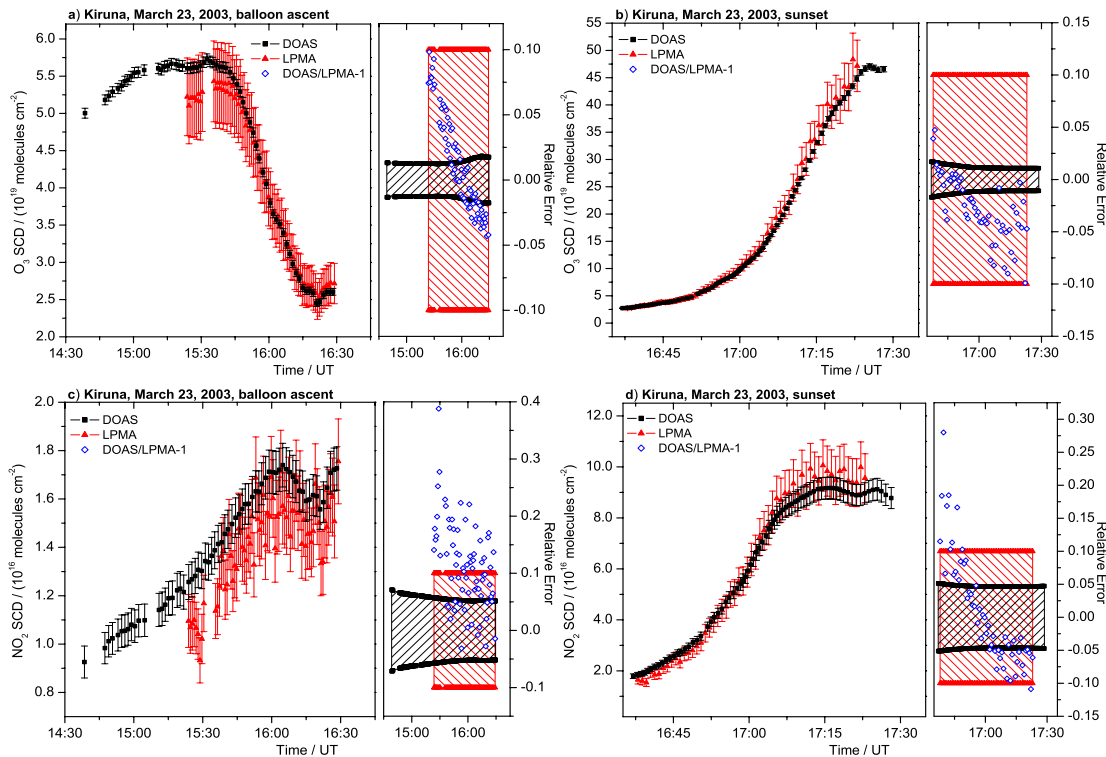


Fig. 1. Slant Column Densities of O₃ (panels **a** and **b**) and NO₂ (panels **c** and **d**) retrieved from LPMA (red triangles) and DOAS (black boxes) direct solar spectra measured at Kiruna on 23 March 2003, during balloon ascent (panels **a** and **c**) and sunset (panels **b** and **d**). Each panel consists of two sub-panels showing the measured SCDs on the left and the corresponding relative errors and deviations (blue open stars), i.e. $\text{SCD}(\text{DOAS})/\text{SCD}(\text{LPMA}) - 1$, on the right. For clarity not all DOAS data points are shown.

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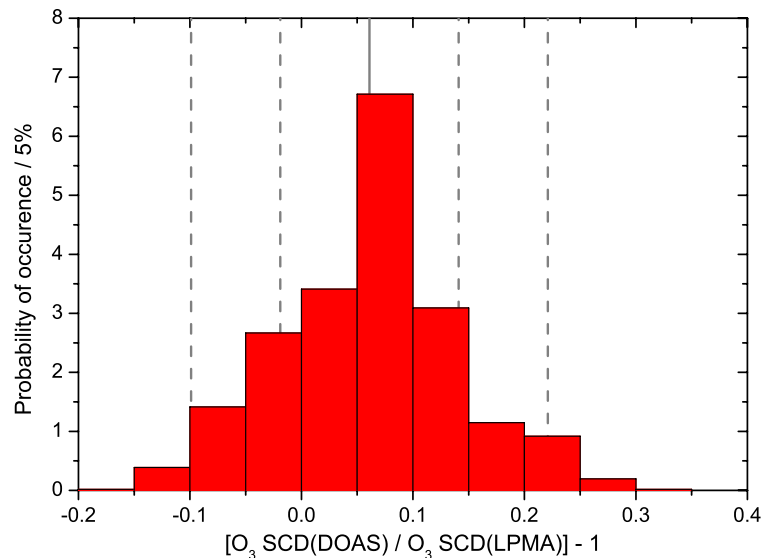


Fig. 2. Histogram of the relative deviations between O₃ SCDs retrieved by DOAS and LPMA, i.e. $\text{SCD}(\text{DOAS})/\text{SCD}(\text{LPMA}) - 1$. The 1032 data points are grouped and counted for occurrence in bins of 5% relative deviation. The mean and the corresponding standard deviation are 6.1% and 8.0%, respectively. The mean value is shown as grey solid line, grey dashed lines correspond to one and two times standard deviation from the mean.

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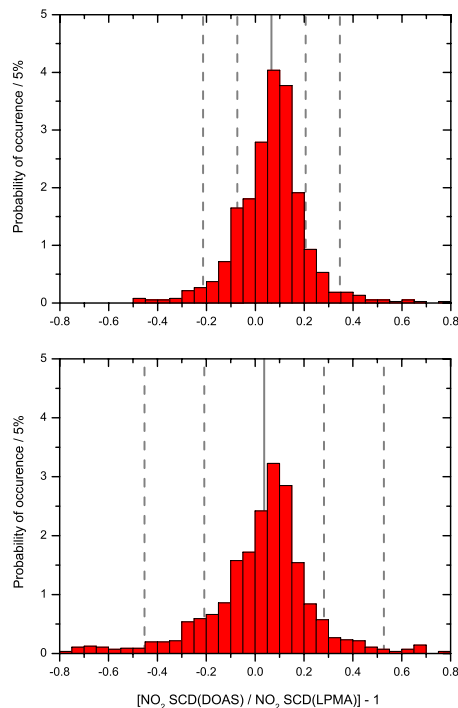


Fig. 3. Histogram of the relative deviations between NO₂ SCDs retrieved by DOAS and LPMA, i.e. $\text{SCD}(\text{DOAS}) / \text{SCD}(\text{LPMA}) - 1$. Lower panel: All 1016 data points are grouped and counted for occurrence in bins of 5% relative deviation. The mean and the corresponding standard deviation are 3.7% and 24.5%, respectively. Upper panel: All data where the corresponding SCDs exhibit relative errors larger than 25% are removed. The remaining 753 data are grouped and counted for occurrence in bins of 5% relative deviation. The mean and the corresponding standard deviation are 6.6% and 14.0%, respectively. Common: The mean value is shown as grey solid line, grey dashed lines correspond to one and two times standard deviation from the mean.

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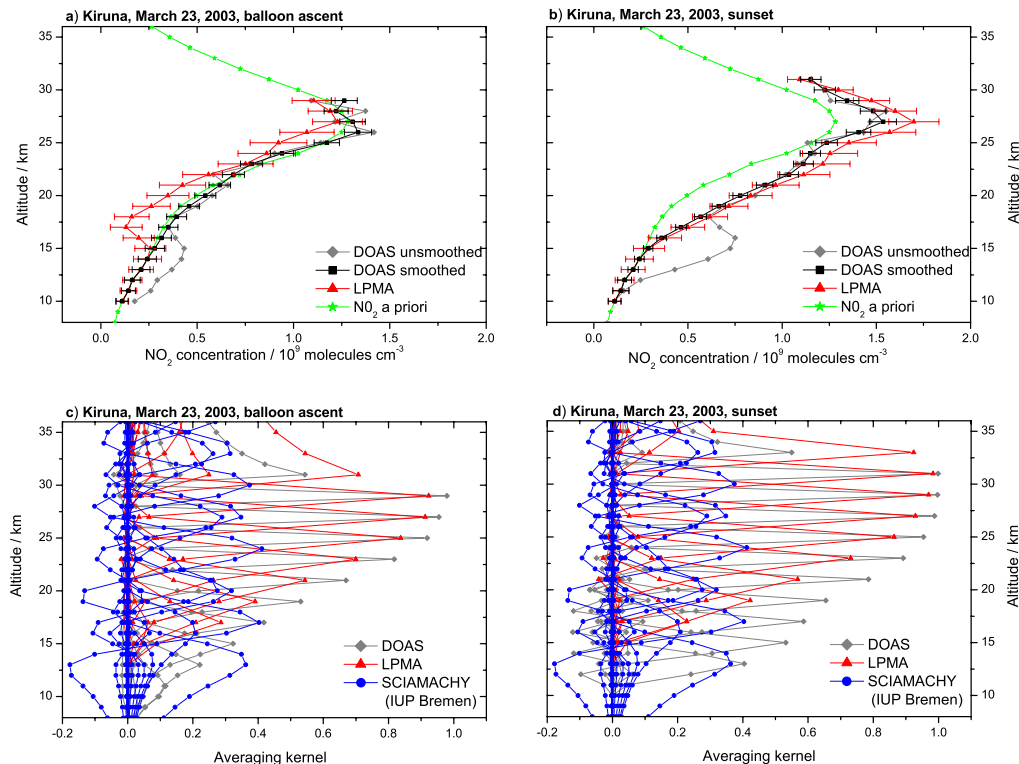


Fig. 4. Vertical NO₂ profiles (panel **a** and **b**) and corresponding averaging kernels (panel **c** and **d**) retrieved from balloon ascent (panel **a** and **c**) and sunset (panel **b** and **d**) for the flight from Kiruna on 23 March 2003. LPMA data (red triangles) are retrieved on a 2 km, DOAS data (grey diamonds) on a 1 km altitude grid. In addition smoothed DOAS data are shown (black boxes). The green stars represent the NO₂ a priori profile. The averaging kernels of the SCIAMACHY NO₂ retrieval are plotted (blue circles) for comparison.

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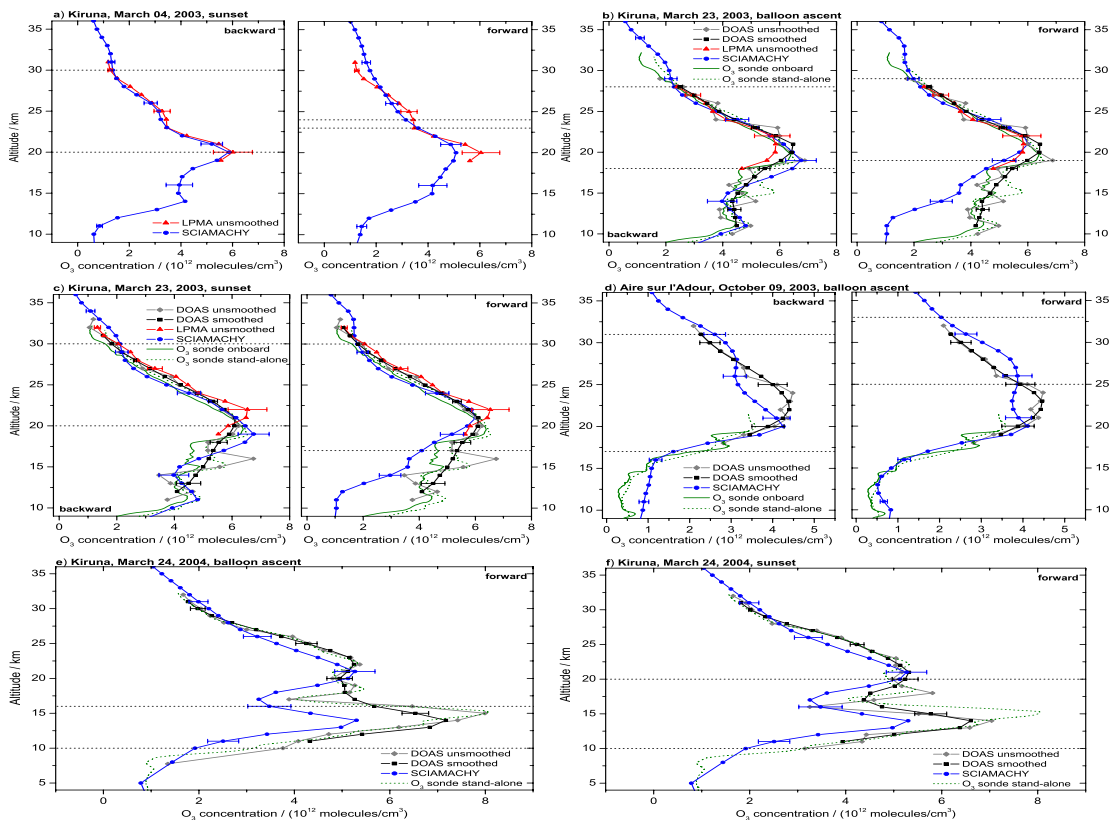


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Fig. 5. Comparison of O₃ profiles inferred from SCIAMACHY limb observations with correlative balloon borne measurements. Panels (a) to (f) correspond to observations at (a) Kiruna on 4 March 2003, at Kiruna on 23 March 2003, during (b) balloon ascent and (c) sunset, at (d) Aire sur l'Adour on 9 October 2003 and at Kiruna on 24 March 2004, during (e) balloon ascent and (f) sunset. The left and right sub-panels correspond to backward and forward coincidences. Satellite data are shown as blue circles. Smoothed DOAS data are plotted as black boxes, LPMA data as red triangles. The grey diamonds represent DOAS O₃ profiles at full altitude resolution without smoothing. If available, O₃ profiles measured by an electrochemical cell onboard the gondola during balloon ascent are shown as green line. The dashed green lines correspond to O₃ profiles inferred from stand-alone in-situ sondes launched from the ground station shortly after balloon launch. The altitude range between the horizontal dotted lines represents the range where coincident air masses are found. Unfortunately on 24 March 2004, only forward match data are available for SCIAMACHY. For better visibility, only selected error bars are shown.

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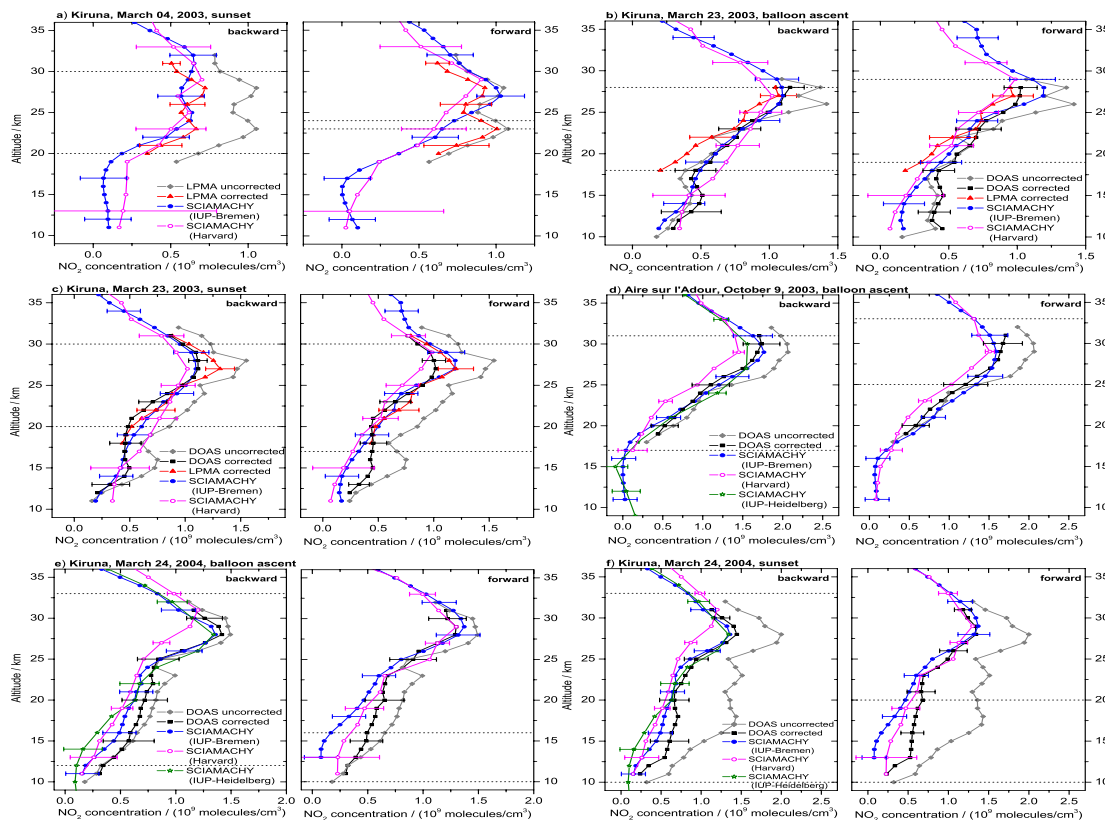


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Fig. 6. Comparison of NO₂ profiles inferred from SCIAMACHY limb observations with correlative balloon borne measurements. Panels **(a)** to **(f)** correspond to observations at (a) Kiruna on 4 March 2003, at Kiruna on 23 March 2003, during (b) balloon ascent and (c) sunset, at (d) Aire sur l'Adour on 9 October 2003 and at Kiruna on 24 March 2004, during (e) balloon ascent and (f) sunset. The left and right sub-panels correspond to backward and forward coincidences. Satellite data inferred by IUP-Bremen are shown as blue full circles, inferred by Harvard as open magenta circles, inferred by IUP-Heidelberg as green open stars. Photochemically corrected and smoothed DOAS data are plotted as black boxes, photochemically corrected LPMA data as red triangles. The grey diamonds represent balloon borne profiles at full altitude resolution without photochemical correction which are taken from the DOAS data except for the flight from Kiruna on 4 March 2003, where LPMA data are used. The altitude range between the horizontal, dotted lines represents the range where coincident air masses are found. For better visibility, only selected error bars are shown.

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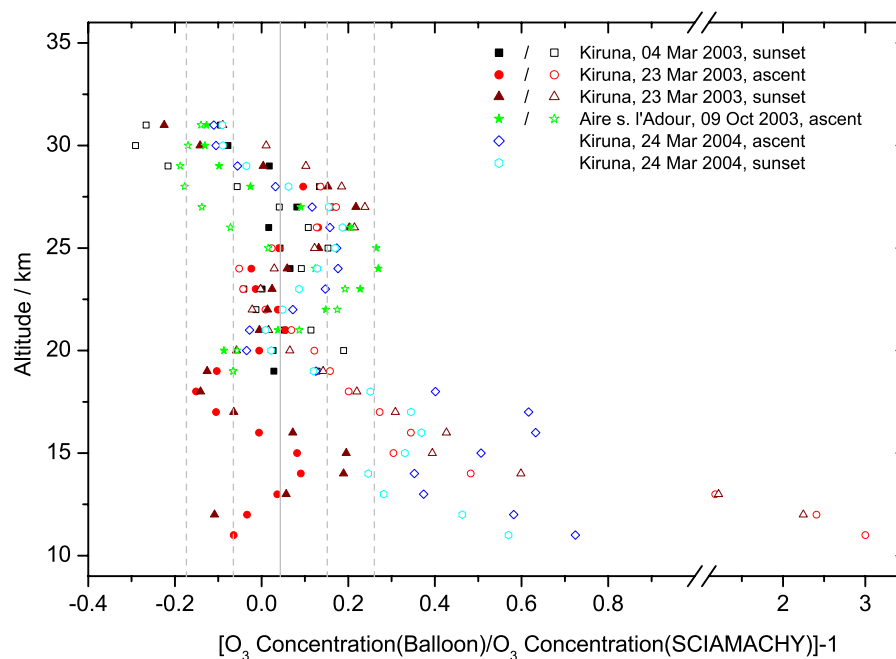


Fig. 7. Relative deviations between satellite and balloon borne measurements of O₃ profiles. Filled and open symbols correspond to backward and forward coincidences, respectively. Observation sites and conditions are indicated in the legend. The mean deviation of all coincident data in the 20 km–31 km altitude range is 4.3% with 10.8% standard deviation. The grey lines indicate the mean at +0.043 and the one and two times standard deviation boundaries with respect to the 20 km–31 km altitude range. Note the broken abscissa.

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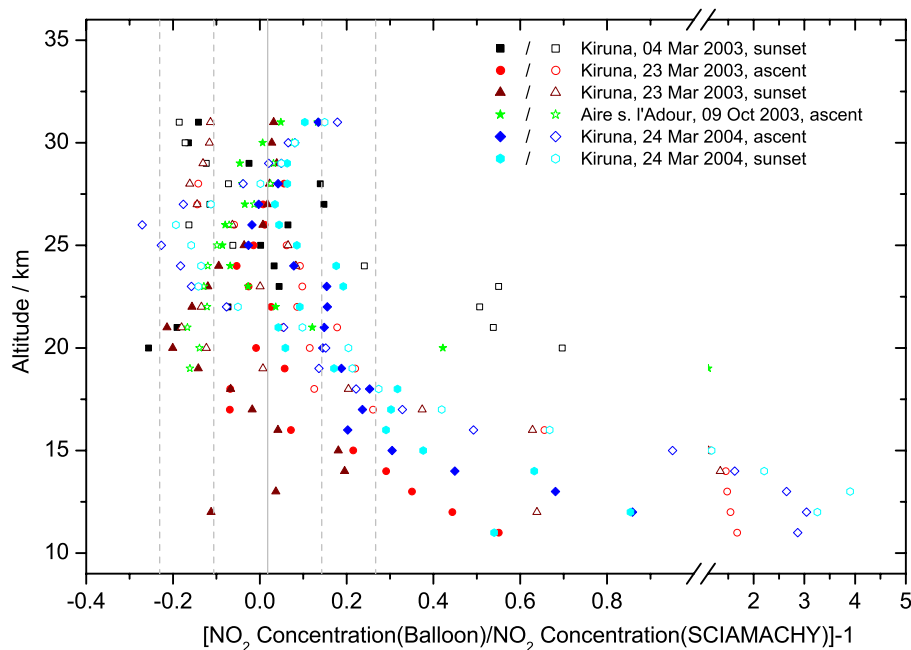


Fig. 8. Relative deviations between satellite (IUP-Bremen) and balloon borne measurements of NO₂ profiles. Filled and open symbols correspond to backward and forward coincidences, respectively. Observation sites and conditions are indicated in the legend. SCIAMACHY data corresponding to the backward coincidence on 4 March 2003, at Kiruna are shifted by -2 km. The mean deviation of all coincident data in the 20 km–31 km altitude is 1.8% with 12.4% standard deviation. The grey lines indicate the mean at 0.018 and the one and two times standard deviation boundaries with respect to the 20 km–31 km altitude range. Note the broken abscissa.

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