

# Intercomparison of satellite retrieved aerosol optical depth over ocean during the period September 1997 to December 2000

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Abstract

Monthly mean aerosol optical depth (AOD) over ocean is compared from a total of 9 aerosol retrievals during a 40 months period. Comparisons of AOD have been made both for the entire period and sub periods. We identify regions where there is large disagreement and good agreement between the aerosol satellite retrievals. Significant differences in AOD have been identified in most of the oceanic regions. Several analyses are performed including spatial correlation between the retrievals as well as comparison with AERONET data. During the 40 months period studied there have been several major aerosol field campaigns as well as events of high aerosol content. It is studied how the aerosol retrievals compare during such circumstances. The differences found in this study are larger than found in a previous study where 5 aerosol retrievals over an 8 months period were compared. However, results in coastal regions are promising especially for aerosol retrievals from satellite instruments dedicated for aerosol research. In depth analyses explaining the differences between AOD obtained in different retrievals are clearly needed. We limit this study to identify differences and similarities and indicate possible sources that affect the quality of the retrievals. This is a necessary first step towards understanding the differences and improving the retrievals.

1. Introduction

Satellite retrievals of aerosols and clouds have given much insight into the problem of quantification of the direct and indirect aerosol effects (e.g. Husar et al., 1997; Kaufman and Fraser, 1997; Nakajima and Higurashi, 1998; Boucher and Tanre, 2000; Nakajima et al., 2001; Tanre et al., 2001; Rosenfeld, 2000; Rosenfeld et al., 2002; Koren et al., 2004). However, significant uncertainties remain regarding the radiative and climate effect of aerosols of anthropogenic origin (Haywood and Boucher, 2000; IPCC, 2001; Ramanathan et al., 2001; Kaufman et al., 2002a). For the direct aerosol effect un-

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certainties exist both due to limited information of spatial and temporal variation in the aerosol optical properties and the composition of the aerosols. Of particular importance is the fact that the crucial parameter single scattering albedo is poorly quantified. Satellite data have greatly improved the knowledge about the distribution of aerosols in the atmosphere. Given the complicated task of retrieving aerosol information from satellite instruments (King et al., 1999), it was perhaps not surprising that Myhre et al. (2004) showed, by comparing 5 satellite aerosol retrievals over ocean for an eight month period (November 1996 to June 1997), that substantial differences in aerosol optical depth (AOD) are present. In general, they found differences in AOD of a factor of two between the different datasets, but in some regions it was even higher. The best agreement in AOD was found in coastal regions with high AOD, whereas the largest discrepancies were found over large areas of remote oceanic regions in the southern hemisphere. Cloud screening was implicated as probably one of the main reasons for the large disagreement.

In this study we investigate AOD over ocean from several satellite aerosol retrievals over a 40 months period from September 1997 until December 2000. This is a much longer period than studied in Myhre et al. (2004) and allows investigation of inter-annual variability in AOD. For this period, 4 different aerosol satellite retrievals are investigated that were producing data for the entire period. Out of these 4 retrievals, 3 were also used in the intercomparison study in Myhre et al. (2004). In addition we focus on two shorter time periods; (i) an 8 months period with one additional satellite aerosol retrieval and two supplementary versions of one of the four main retrievals, (ii) a 10 months period with two additional retrievals for dedicated aerosol research. A particularly interesting issue is to see how the long term monitoring satellite retrievals compare to retrievals from satellite instruments dedicated (e.g. such as POLDER, MODIS, MISR) to investigation of aerosols.

A significant advantage of our intercomparison of AOD in this 3 years period compared to the earlier intercomparison period is that much more ground based sunphotometer data from AERONET are available. This allows a broader comparison between

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the satellite aerosol retrievals and the AERONET measurements and furthermore an evaluation of under which conditions differences in the retrievals are largest. The aim of this study is to provide data for the global modelling community for comparing and improving global aerosol models. As our comparison of satellite data with AERONET data is made in this context, it is not made on a spatial and temporal resolution which is high enough to make a detailed study of the deficiencies of the satellite retrievals. Finally, we also compare AOD from the various satellite retrievals in some selected regions and time periods with particular focus on e.g. episodes of large AODs or measurement campaigns. Overall, we explore whether differences in AOD are particularly large for e.g. certain satellite retrievals, oceanic regions, aerosol sizes, and ranges of AOD.

2. Method

For the comparison of satellite retrievals, data for the time period September 1997 to December 2000 have been collected. During this entire period, continuous AOD data from four satellite retrievals are available, namely AVHRR, (a one (AVHRR-1) and a two channel (AVHRR-2) retrieval), TOMS, and SeaWiFS. In addition AOD data from VIRS for an 8-month period (January to August 1998) are available. Further, two versions of the one channel AVHRR retrieval for the same time period are available. Since March 2000 data from the retrievals of the dedicated aerosol instruments MODIS and MISR (both onboard the Terra satellite) have become available, and thus we include 10 months of data for these retrievals. For MODIS version 4 data have been used. Due to the MISR capacity of multi angle viewing there is a rather small swath width for this instrument, requiring 9 days for full global coverage. Thus monthly mean datasets, which are used in much of this study, contain fewer data points in time for MISR than for the other retrievals. From September 1997 to August 1998 data for some limited regional areas (Meteosat Second Generation field of view) are included for a retrieval based on a combination of GOME and ATSR-2. The GOME/ATSR-2 retrieval yields only 3

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days of measurements each month due to pixel size mode programming. Therefore monthly averages need to be compared with care. It should be noted that this retrieval method was mainly developed for application over land and separation of the basic aerosol components. Due to the large GOME pixel size of  $80 \times 40 \text{ km}^2$  retrieval values at coastal sites may be dominated by AOD values derived from dark surface reflectances over land. AVHRR has been producing data since 1981 and TOMS, except for a break from 1992 to 1996, since 1979. September 1997 was chosen as the first month of the comparison because this was the first month with AOD data from SeaWiFS. Currently, since the SeaWiFS data processing is optimized for the ocean color measurements; very thick aerosol AOD cases such as the dust and smoke plumes are usually masked out due to large uncertainties in the ocean color products in these cases. SeaWiFS has a reflectance threshold at 865 nm corresponding to AOD of  $\sim 0.3$  above which aerosol retrievals are discarded. Thus, the SeaWiFS AOD is mostly applicable and valid in the open ocean regions (Wang et al., 2000a, b). It should be noted that AVHRR-2 has a similar AOD threshold of 1.0 at 550 nm. Table 1 gives a short description of the 8 retrievals.

The AOD data are reported for different wavelengths; AVHRR-1 at 630 nm, AVHRR-2 at 550 nm, SeaWiFS at 865 nm, TOMS at 380 nm, VIRS at 630 nm, GOME/ATSR-2 at 550 nm, MODIS at 550 nm, and MISR at 558 nm. To make the most meaningful comparison possible, we convert the AOD data for the different wavelengths to the AOD at 550 nm. For SeaWiFS the Ångström coefficient reported as part of the SeaWiFS data product is used for conversion. The AVHRR-1 data and the VIRS data have been converted using the Ångström coefficient from AVHRR-2. For TOMS a provisional estimate of AOD at 550 nm is made by scaling the 380 nm values with the monthly mean ratio between 550 nm and 380 nm AOD values from the period 1979 to 1992 for each grid point. For AVHRR-2, MODIS, GOME/ATSR-2 and MISR the reported wavelength are used. All the satellite data are compared for a  $1 \times 1^\circ$  grid on a monthly mean basis.

The satellite retrievals are compared with ground based AOD measurements from

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AERONET (Holben et al., 1998). Because the satellite retrievals in this comparison mostly include data over ocean only, measurements from AERONET stations located on islands and near coastlines are used. Even so, the grid square where the station is located will in many cases be predominantly over land and therefore have no AOD value from the satellite retrievals. The AERONET data are therefore compared to an average of 9 (3 in latitudinal and 3 in longitudinal direction)  $1 \times 1^\circ$  grid squares of satellite data, where the ground station is located in the centre grid square. Note that the comparison of satellite data to the AERONET data is similar to a model to AERONET comparison, since currently horizontal resolution is around  $3^\circ$  in many global models. Based on the MODIS data, which are also retrieved over land, the assumption of using 9 grid points instead of 1 in the comparison with AERONET is investigated. The results of this investigation is presented in Appendix A, showing generally very similar results for 9 grid points compared to the closest grid point to an AERONET station. Out of 33 AERONET stations used in this study, this assumption seems unsatisfactory for only 3 stations.

Remote sensing from space and ground based sunphotometers use different techniques in retrieval of aerosol information. Satellite retrievals use reflected sunlight at the top of the atmosphere, whereas sunphotometers use direct sunlight transmitted through the atmosphere.

AOD from sunphotometers are usually used as reference since it is based on fewer assumptions and has been shown to give accurate results compared to other measurements (e.g. Haywood et al., 2003a). However, due to particle size and shape distribution and potentially other factors results from the two techniques may not be matching to better than a certain level. Investigations should be made to find this precision level.

### 3. Intercomparison of satellite AOD

Figure 1a shows the geographical distribution of AOD for a mean of 40 months for 4 satellite retrievals. The figure confirms results of many previous satellite studies that

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AOD is generally highest near continental regions. This is evident in all four data sets. The aerosol plume over the Atlantic west of Africa is clearly evident, with mineral dust dominating in the northern region and biomass burning dominating in the south and a mixture of these aerosols around Equator. As in Myhre et al. (2004) significant differences can be seen in the southern hemisphere at high latitudes. SeaWiFS has generally lower AOD than the three other retrievals. Figures 1b and c show the geographical distribution of AOD for the 2 sub periods. The VIRS data show a similar pattern in AOD as the 4 retrievals which are available for the whole period in this study. The magnitude in AOD for VIRS is most similar to AVHRR-1 and TOMS. GOME/ATSR-2 has a sparse geographical coverage of AOD but indicates gradients and AOD values west of Africa that are similar to the other aerosol retrievals. The improved quality control in AOD for AVHRR-1 reduces the coverage of AOD, but influences the magnitude of AOD only weakly where data are available.

MODIS and MISR have higher AOD values in many regions compared to the 4 other retrievals shown in Fig. 1c, with MISR AOD usually somewhat higher than MODIS values. The highest AOD values are off the coast of northwest Africa and at high latitudes. The magnitude in AOD varies somewhat between the various retrievals but the pattern in AOD in the tropical Atlantic Ocean is rather robust. The limited coverage of AOD for some of the retrievals is due to orbital drift. It is interesting to see the values up to around 0.3 in the southern hemisphere and up to 0.6 in the northern hemisphere at high latitudes for MODIS and MISR. In the region in the southern hemisphere the aerosols are mainly sea salt and to some extent natural sulfate aerosols and episodic anthropogenic aerosols. This is a region with few AERONET measurements and has been identified in Myhre et al. (2004) as the region with largest difference between the various satellite retrievals. The high AOD in the northern Pacific is interesting and is a likely a combination of outflow of aerosols from the Asian continent and sea salt aerosols generated by the relatively high windspeeds in these regions. Note also the elevated AOD around 10–20° N in the eastern part of the Pacific Ocean which is seen to various extent in all the datasets. Two areas with particular differences are near the

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coast of China and the North Sea. In the former region the AOD varies gradually from about 0.2 to 1.0. In the North Sea MISR shows a particularly high AOD.

In Fig. 2 the global mean of all the 8 data sets with global coverage are shown. For Fig. 2a the whole period is displayed whereas in Fig. 2b the 8 months sub-period is highlighted. The global mean AOD differs by at least a factor of 2. AVHRR-2 and SeaWiFS have considerably weaker annual variability than the other data sets. The inter-annual variability differs somewhat between the AVHRR-1 and TOMS. The AOD from the VIRS retrieval is in the upper range of the AODs represented in the study. A variation over the 8 months can be seen in the VIRS data. Interestingly, a comparison between the 2nd generation and 3rd generation AVHRR-1 product yields differences that are almost as large as differences between various aerosol retrievals on different satellite platforms. The two AVHRR products have similar temporal variation in AOD over the 8 months period. Figure 2c highlights the second sub period indicating weak variation during the 10 months period for the retrievals, except for AVHRR-1 (partly AVHRR-2) due to orbital degradation. The yearly variation in global AOD is very similar for MODIS and MISR despite differences in their magnitude.

Figure 3 shows the zonal mean AOD for the entire 40 month period (a), the 8 months period from January to August 1998 (b), and the 10 months period from March to December 2000. Similar to what was found in Myhre et al. (2004) the largest differences are found at high latitudes, whereas the results in tropical regions are generally more comparable. The largest differences between MODIS and MISR in the monthly mean are found at high latitudes, in particular in the southern hemisphere. For the entire period the differences are largest in the southern hemisphere, whereas in the latest sub period large differences are also found in the northern hemisphere. A common feature with local maxima near equator and at higher latitudes and local minima at 20–30° in both hemispheres can be seen to various extent in all the retrievals.

The aim of Fig. 4 is to illustrate the annual cycle of AOD, and its inter-annual variability, over large oceanic regions. The thick lines show the 40 months period mean AOD and the dotted lines show minimum and maximum AOD values during the same pe-

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riod. The oceanic regions are defined in Appendix B. AVHRR-1 and TOMS have much larger inter-annual variability than AVHRR-2 and SeaWiFS. This is also indicated in Fig. 2a, but Fig. 4 shows that this is the case over all the 5 oceanic regions. Generally, AVHRR-1 and TOMS have a larger inter-annual variability from April to November than during the rest of the year.

Figure 5 shows AOD from the nine data sets for 11 smaller oceanic regions, which are in coastal areas, attempting to cover the oceanic regions with highest AOD. AOD is thus usually higher in these regions than the global mean. Further, in general the difference between the AOD from the various retrievals is smaller than for global averages and for larger oceanic regions, as found also in Myhre et al. (2004). In about half of the 11 regions a distinct seasonal variation in AOD can be seen. However, the seasonal amplitude varies between the satellite retrievals. For many of the retrievals the agreement is best at the East coast of USA, and over the Arabian Sea and the Red Sea. The differences between the two AVHRR-1 data sets are smaller than shown in Fig. 2 for global conditions. The VIRS data have many similarities with the AVHRR-1 data sets, although often with somewhat higher values. The agreement in AOD between MODIS and MISR is very good, except for the Caspian Sea where the difference between the retrievals is usually large. Further, there is a tendency that MODIS and MISR as well as GOME/ATSR-2 have higher AOD than the other retrievals.

In Fig. 5 an average value for each region was shown. However, despite two datasets having a similar average value, they may have a different spatial distribution. Figure 6 shows the spatial correlation coefficient between various retrievals for the months with common availability of data. The correlation coefficient between the satellite retrievals is shown for the oceanic regions illustrated in Fig. 5 for six of the retrievals. The highest correlation coefficient is found for the Angola basin, the Red Sea, and the Arabian Sea, whereas the weakest correlation coefficient is identified in the Caspian Sea. Also in the Mediterranean Sea and the Black Sea the correlation coefficients are low. Two apparent explanations of the variations in the correlation coefficient for the regions are differences in the temporal variation in AOD and variation in aerosol type and mixture.

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5 The Mediterranean Sea and Caspian Sea is likely to be dominated by aerosols from industrial pollution with episodic influence by mineral dust plumes. This yields significant temporal variation in AOD. Furthermore, and probably more important is the fact that many aerosol types and sizes are likely present, sometimes even in complicated internal mixtures. On the other hand the Angola basin and the Red Sea are mainly dominated by one aerosol type (besides some sea salt aerosols); namely biomass burning aerosols and mineral dust, respectively. For AVHRR-1 the correlation with other aerosol retrievals is usually high, not surprisingly it is generally best with VIRS but also highly correlated with MODIS and MISR. For AVHRR-2, TOMS, and SeaWiFS 10 the correlation with other retrievals is usually slightly lower than for AVHRR-1. The correlation between these three retrievals is low, and each of them are normally more strongly correlated to the four other retrievals. The spatial correlation between MODIS and MISR is high, with exceptions for the East coast of USA, the Black Sea, and especially for the Caspian Sea. However, it is interesting that AVHRR-1 often has even higher correlation with both MODIS and MISR than the internal correlation between MODIS and MISR. The internal correlation between MODIS and MISR is certainly influenced by the low temporal sampling of MISR as discussed above. Note here that in some regions the analysis with AVHRR-1 against MODIS and MISR is based on relatively few months due to data availability.

20 The spatial and temporal variation in the aerosol distribution is large, and in some periods episodes of large aerosol amount can occur. This can either be due to large fires (of natural or anthropogenic origin), occurrences of large mineral dust outbreaks or when meteorological conditions favor high concentration of aerosols resulting from industrial activity. The aim of Fig. 7 is to illustrate how the various satellite retrievals compare under such circumstances. We have selected some episodes during the period of this study with high amounts of aerosols. Aerosol campaigns have taken place in many regions around the world, often focusing on areas with large AOD. In Fig. 7 we have included data from regions close to three aerosol campaigns (INDOEX, 25 SAFARI 2000, and SHADE). The agreement between the satellite retrievals is best

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when AOD is generally low, and largest differences are found when AOD is high. In particular, the large biomass burning events in Indonesia in October 1997 and Mexico in May 1998 reveal the largest differences between the satellite retrievals. Actually, the differences can be as large as a factor 3–4. In regions dominated by industrial pollution the results are generally similar or in better agreement than for global mean conditions. One explanation for the difference in the monthly mean AOD during events of high AOD is different sampling. This is related to different swath width and time of overpass that influence the cloud screening and glint mask. Therefore, we can expect that sampling issues are more evident in a small region with high variability than in the large temporal or spatial averages. The relative differences in AOD for the retrievals were greater for SAFARI 2000 than for INDOEX and SHADE, even when considering only the four retrievals with observations in all campaigns. The AOD was also larger for the region close to SAFARI 2000 than for the INDOEX and SHADE regions. AVHRR-2 and SeaWiFS have usually the lowest values, most likely linked to the upper threshold values for AOD. For the cases including MODIS and MISR data these two retrievals have high AOD, and the agreement between them is good and generally better than the agreement between any other two retrievals.

#### 4. Comparison of AOD between aerosol satellite retrievals and AERONET

A large advantage compared to Myhre et al. (2004) is that in this study a much larger set of AERONET observations are available for comparison with the satellite retrievals. This is both due to a longer time period of investigation and to the fact that more AERONET stations were in operation. Validations of the individual aerosol satellite retrievals have been performed for several of the satellite products (e.g. Stowe et al., 2002; Torres et al., 2002; Remer et al., 2002; Holzer-Popp et al., 2002b; Kahn et al., submitted, 2004<sup>1</sup>; Liu et al., 2004) Comparison between the satellite retrievals and

<sup>1</sup>Kahn, R. A., Gaitley, B. J., Martonchik, J. V., Diner, D. J., Crean, K. A., and Holben, B.: MISR global aerosol optical depth validation based on two years of coincident AERONET ob-

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the AERONET as a function of time is shown Fig. 8 for 33 stations. In general a good agreement is found between the AERONET and satellite retrievals, and seasonal and inter-annual variations seen in the AERONET data are mainly captured also in the satellite data. The best agreement between the AERONET data and the satellite retrievals are probably found at Wallops, Barbados, and Bermuda, whereas at Dakar significant differences are found in the first part of the period where we use the less validated level 1.5 data.

In Fig. 9 a subset of the AERONET stations with longest time series of observations are further studied. Mean and standard deviation for the AERONET data and satellite retrievals are given in the figure, as well as the corresponding correlation coefficient for each station. No obvious systematic differences in the satellite retrievals in their comparison with the AERONET averages are apparent. At Lanai and San Nicolas the mean AOD is higher in all the satellite retrievals compared to the AERONET data. Otherwise no systematic differences can be found for the mean AOD. The standard deviation is smaller or equal in the AERONET data than in the satellite retrievals at Lanai and San Nicolas. SeaWiFS has in often the lowest standard deviations of the seven data sets. The correlation coefficient is usually high and there are no specific AERONET stations with a consistently lower correlation coefficient, but for many of the satellite retrievals it is low at San Nicolas. Bermuda, Bahrain, and Wallops are stations with the generally highest correlation coefficient.

Note here that we use a rather large spatial resolution as described in Appendix A and the spatial variation in AOD can be significant. In particular variation in relative humidity can influence the spatial variation (Haywood et al., 1997; Myhre et al., 2002; Anderson et al., 2003) as water uptake increases the AOD for hygroscopic aerosols considerably. AERONET measurements are taken regularly when clouds are not present. Satellites measure over all regions and thereafter a cloud screening is performed and will thus generally include more aerosol information in regions close to cloudy regions. Relative humidity is frequently higher close clouds than in clear sky, and thus AOD may

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be higher closer to clouds. On the other hand, clouds also wash out aerosols, reducing the AOD close to clouds. Koch et al. (2003) found that over land in Europe and North America that clouds and sulfate aerosols are anticorrelated, indicating that clouds are more efficient in washing out sulfate aerosols than in contributing to aqueous-phase production of sulfate aerosols. However, at high relative humidity the AOD increases significantly due to uptake of water, is therefore not unlikely that AOD is often higher close to clouds than in clear sky. This is in fact what is found based on AERONET and cloud observations in Kaufman et al. (2002b) as well as in Ignatov and Nalli (2002) and Ignatov et al. (2004a). However, to which extent AOD is higher close to clouds than in clear sky regions should be investigated in several other regions, since there could be large regional variations and differences between cloud types and aerosol types.

Figure 10 shows scatter plots of the satellite retrievals against AERONET data for all the 33 stations included in this study. Note that for the eight retrievals shown in this figure the time period varies for which analysis are performed, and that we have removed level 1.5 data from the analysis in the scatter plot analysis. There is a tendency that the satellite retrievals have a weak overestimation compared to AERONET for low AOD, but often underestimate high AOD from AERONET. This pattern varies significantly between the various retrievals. SeaWiFS has a good agreement with the AERONET data for low AOD, but the largest underestimation among the satellite data sets for high AERONET AOD values. TOMS has a very few high AOD values that are clear overestimates compared to the AERONET data. AVHRR-1, AVHRR-2, and VIRS differ also most compared to AERONET for high AOD values. MODIS and MISR have relatively more high AOD values than AERONET values compared to the other retrievals. However, these scatter plots reveal that particularly MODIS and partly MISR generally compare better to AERONET than the rest of the retrievals. Also AVHRR-1 has a slope indicating a good agreement with AERONET given all the uncertainties in this comparison which is based on monthly mean data. The two versions of AVHRR-1 differ significantly, due to different sampling and retrieval procedures used. Two issues regarding this analysis need to be emphasized. The AERONET data used in Fig. 10

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are not the same for all retrievals due to differences in time periods. Considering only data only in the 10 months period in 2000 impacts the results for AVHRR-1, AVHRR-2, TOMS, and SeaWiFS compared to AERONET data to a limited extent. On the other hand including level 1.5 in this scatter analysis would reduce the agreement for these four retrievals with the AERONET data, in particular the underestimation of high AOD values.

When considering monthly means in the comparison of satellite retrievals with AERONET, data sampling issues are important. For the AERONET measurements high frequency data are averaged to daily means. However, the number of days that are included in the AERONET monthly mean of AOD is highly variable. Especially three AERONET monthly data points differ from the satellite retrievals, and they are all based on less than three days of measurements during the whole month. These are Dakar (1 day in August 2000), Taiwan (2 days in October 2000), and Rame Head (3 days in September 1998). Removing these data points from the analysis increased the slope of the regression-line for all satellite retrievals, except for VIRS and AVHRR-1(3g/QC). For the other retrievals the slope of the regression-line changed from 0.04 to 0.26, with the largest change for MODIS and thereafter MISR. Further, removing all AERONET data with less than three days of measurements during the whole month (30 out of a total of 488 data points) had a relatively small impact, less than 0.05 change in the slope of the regression-line compared to the case when only the three data points discussed above were removed. In most the cases the monthly mean data are based on more than 10 days of measurements. This is the case in 375 of the 488 AERONET monthly mean data used in this study.

For AVHRR-1(2g) a large difference from AERONET occurs mainly at Bahrain, in addition to the cases discussed above. However, removing the data points from Bahrain in the analysis influences the slope of the regression-line only to a small extent. For AVHRR-2, TOMS, and SeaWiFS there are no stations with particularly large differences compared to AERONET. For MISR a few AERONET stations reduce the slope of the regression-line significantly. These are in particular Helgoland and Gotland, and

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to some extent Taiwan (October 2000 as mentioned above). Removing these data from the analysis for MISR resulted in a regression-line slope of around 0.85 (also shown in Fig. 10), which is a typical value found in the global MISR validation paper by Kahn et al. (submitted, 2004)<sup>1</sup>. The high MISR values at Helgoland and Gotland for MISR can be seen in Figs. 8 and 1c. The high values found in the North Sea for MISR are probably related to cloud screening and to the use of a climatology of the near-surface wind speed for the ocean white cap model. This may yield too high AOD under high wind events under cloud-free conditions (Kahn et al., submitted, 2004<sup>1</sup>). Removing the same values in the analysis for MODIS as for MISR increased the slope of the regression-line for MODIS similar to MISR. Whereas, our comparison with AERONET is based on monthly mean data, the study of Abdou et al. (submitted, 2004)<sup>2</sup> compare coincident MISR and MODIS AOD with AERONET data during 3 month in 2002 with MISR showing an agreement with AERONET which is at least as good as for MODIS.

In Fig. 11 the scatter plots of AERONET data and the satellite retrievals are divided into three groups according to the AERONET Ångström exponent to identify whether differences are related to the size of the aerosols, which is indicative of aerosol type. For AVHRR-1 and TOMS the results seem rather independent of the Ångström exponent. For AVHRR-2 the results indicate reduced agreement with the AERONET data for the smallest aerosols (high Ångström exponent). In the case of the SeaWiFS data, Fig. 10 indicates an underestimation of large AOD from AERONET. Figure 11 indicates that this is least pronounced for the smaller particles. For MODIS there is no systematic difference with Ångström exponent, whereas for MISR the agreement is best for the largest particles (small Ångström exponents). The values discussed above for MISR at Helgoland and Gotland, and Taiwan influence the analysis for the small particles.

<sup>2</sup>Abdou, W. A., Diner, D. J., Martonchik, J. V., Bruegge, C. J., Kahn, R. A., Gaitley, B. J., Crean, K. A., Remer, L. A., and Holben, B.: Comparison of coincident MISR and MODIS aerosol optical depths over land and oceans scenes containing AERONET sites, J. Geophys. Res., submitted, 2004.

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5. Summary and discussion

In this study monthly mean aerosol optical depth (AOD) is compared from a total of nine aerosol retrievals during a 40 months period, from September 1997 to December 2000. We have identified that differences in various satellite retrievals are substantial and even larger than found in an earlier study based on five different aerosol retrievals during a period of eight months prior to the period analysed here. Aerosol remote sensing from space is a complicated task involving a wide range of physical processes that must be taken into account. Issues related to cloud screening are particularly important. It appears that one problem is that, in many retrievals, the cloud screening is not strict enough resulting in AOD being contaminated by clouds. On the other hand it also appears that some aerosol retrievals are too strict, i.e. high aerosol loadings are classified as clouds and thus no aerosol information is retrieved. In this study we have seen examples of aerosol retrievals adopting upper threshold values for AOD in an effort to avoid cloud influence. For small particles (e.g. from industrial pollution or biomass burning) this procedure could be improved by introducing an additional criterion for the Ångström exponent. However, this is more difficult for larger particles (e.g. mineral dust and sea salt) with smaller Ångström exponents more similar to those of clouds. For example, retrieval of aerosol information under major dust episodes, where AOD can be significantly above 1.0 is particularly difficult. To distinguish heavy dust loads from clouds is difficult and multi channel information is needed. Dedicated aerosol satellite instruments have this capacity and therefore this is a tractable problem for these retrievals. Additionally, in conditions of heavy dust loading, sunphotometers may screen out heavy dust loadings by miss-classification as cloud. During the SHADE campaign there was an indication that during the period of maximum AOD during a major dust storm, the procedure for processing level 1.5 to level 2 sunphotometer data led to rejection of much of the sun-photometer data (Haywood et al., 2003b). Overall, it cannot be ruled out that both sunphotometers and satellite retrievals miss-classify some of the major dust storms as clouds and thus are biased towards lower dust

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conditions.

Despite differences in AOD are substantial, there are also many promising results. The agreement with regard to spatial and temporal distribution in AOD between the two dedicated aerosol instruments in many of the subregions investigated in this study is impressive. This finding is both based on the averaged AOD and its variation in magnitude, as well as spatial and temporal correlation coefficient. Furthermore, in several regions the other aerosol retrievals compare well to MODIS and MISR. It seems that for comparisons in smaller regions the agreement between the aerosol retrievals is best where the influence of only few aerosol types is typical.

The analysis performed in this study has been used to identify regions with patterns of agreement and disagreement. The seasonal variation in AOD is well reproduced by the aerosol retrievals at Angola Basin, east coast of USA, Arabian Sea, and Red Sea. The magnitude of the seasonal variation differs between the retrievals, but their timing of maximum in AOD in June or July is very similar. For these four regions the main aerosol components varies substantially such as biomass burning aerosols, aerosols from fossil fuel use, and mineral dust. The spatial correlation in AOD for the aerosol retrievals shows also good results for three of these regions, namely Angola Basin, Arabian Sea, and Red Sea. In addition the spatial correlation at the Cape Verde Plateau is high between many of the retrievals. Several of the AERONET stations show good agreement. The pattern found at the stations Bahrain, Barbados, Bermuda, Venice, and Wallops are the most encouraging.

This study clearly shows that the disagreement between aerosol satellite retrievals is particularly large during events of large influence of aerosols, with differences in AOD over a factor of 3. Part of this difference arises from upper thresholds in AOD, but this can certainly not explain all the difference. During the INDOEX campaign the difference in AOD between four of the retrievals was relatively small, and smaller than during the other major aerosol campaigns. It is noticeable that MODIS and MISR mostly have higher AOD during the events studies here than the other retrievals. We have identified the Caspian Sea as the region having the largest disagreement between the aerosol

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retrievals. This is regard to the magnitude of AOD, its temporal variation, and spatial correlation. The spatial correlation between the aerosol satellite retrievals is very weak for the Caspian Sea and for MODIS and MISR it is even weakly anti-correlated. The Black Sea and Mediterranean Sea are also regions with poor agreement. Particularly large differences between the satellite retrievals are found over remote oceanic regions, in particular at high latitudes southern hemisphere at the edge of possible retrieval of aerosols. The agreement with AERONET data seems particularly poor at Gotland, Helgoland, Rame Head, and Taiwan.

Comparisons with AERONET data reveal differences among the satellite aerosol retrievals, with MODIS data giving generally the best agreement. Also MISR and AVHRR-1 compare very well against the AERONET data. Note here that that the comparison between AERONET and satellite retrievals is based on monthly mean data, so that many factors may influence the comparison such as sampling and cloud screening. A more detailed comparison with AERONET data should be based on daily data, to explore differences and evaluate different aerosol retrievals.

A wide range of factors determine the accuracy of the retrievals and thus how well different retrievals compare. Note here that an agreement in monthly mean AOD in certain regions could be a coincidence and due to compensating errors. Among factors important for the quality of the remote sensing of aerosols are both connected to instrumental designs and to retrieval limitations or weaknesses. Instrumental dependences related to calibration issues, numbers of spectral channels available for aerosol retrieval with weak of overlap with non homogenously distributed gases in the atmosphere and potential for cloud screening, and spatial resolution are crucial elements. Thereafter, there are several factors influencing the quality of the retrieval algorithm such as choice of radiative transfer code, treatment of surface reflectance, cloud screening procedure, and aerosol microphysical model. The potential of the aerosol microphysical model used in the retrieval depends to a large degree on the instrumental designs, but also several user specifications are important. This is for instance related to the single scattering albedo of the aerosols. In our intercomparison study also different equatorial

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passing times could be of importance, despite Kaufman et al. (2000) showed that diurnal variation in AOD was small. However, the equatorial passing time influences the sampling of the data, such as the cloud screening, sun glint screening, and quality control.

5 To explain the causes of the various analyses in this study a very thorough and tedious investigations must be made. First of all detailed analysis based on daily data must be performed, including comparison with ground based sunphotometers, lidars as well as ideally with aircraft measurements. An investigation of how a comparison on daily data relates to a monthly comparison is also needed. To fully understand  
10 the causes of differences between datasets a detailed investigation including testing of various parts of the algorithms needs to be made.

Appendix A

To assess whether the eight nearby grid points can be representative in addition to the grid point with an AERONET station we use MODIS which retrieves data over land and ocean. Note here that we here in many cases apply the ocean retrieval described in  
15 this study as well as a land retrieval (Kaufman et al., 1997). In Fig. A1a a scatter plot between MODIS average (9 grid points) and MODIS centre (1 grid point) is shown for 10 months for all the 33 AERONET stations. Overall the agreement is reasonably good with some few cases where the MODIS centre has somewhat higher AOD than the MODIS average. Figure A1b shows a scatter plot of MODIS data with AERONET data  
20 for both MODIS average and MODIS centre. Results for the two datasets show rather similar agreement with AERONET data. To investigate whether the differences found for MODIS average and MODIS centre in Fig. A1a and A1b are particularly large for certain stations, results for each station are given in Fig. A1c. Figure A1c shows that  
25 the agreement between MODIS average and MODIS centre is very good except for 3 locations, namely Arica, Bermuda, and Nauru. Therefore, using 9 grid points in the comparison with the AERONET data seems reasonable, but some precaution when

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comparing with Arica, Bermuda, and Nauru should be taken. In a detailed analysis of the performance of the satellite retrievals higher resolution data than  $1\times1^\circ$  should be used.

Appendix B

5 In this study we have analyzed data adopting in many cases datasets selected for certain regions and certain time periods. Figure A2 shows the area definition of the oceanic regions chosen in this study. In Fig. A2a the 5 major oceanic regions are given (used in the presentation in Fig. 4), whereas in Fig. A2b the 11 subregions are shown (used in Figs. 5 and 6). The regions for the episodes (as given in Fig. 7) with high  
10 AODs are illustrated in Fig. A2c.

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**Table 1.** Description of aerosol satellite retrievals.

Satellite instrument	AVHRR-1*	AVHRR-2	TOMS	SeaWiFS
Responsible scientists	A. Ignatov	M. Mishchenko and I. Geogdzhayev	O. Torres	M. Wang
References	Stowe et al. (1997, 2002)	Mishchenko et al. (1999, 2003); Geogdzhayev et al. (2002, 2004)	Torres et al. (1998, 2002)	Gordon and Wang (1994)
Aerosol microphysics particles,	Spherical with mono-modal log-normal size distribution ( $R_m=0.1\ \mu\text{m}$ in $dN/dR$ representation; $\Phi=2.03$ ; $n=1.40-0i$ )	Spherical aerosols ( $n=1.5-0.003i$ ) with a power-law size distribution. The refractive index is wavelength independent and includes some aerosol absorption	Three aerosol types: sulfate, carbonaceous and desert dust. Lognormal size distributions. Prescribed real ref. Index for each aerosol type. Imag. ref. index is retrieved.	Shettle and Fenn (1979) aerosol models. Bi-modals log-normal size distributions. Values of single scattering albedo from 0.93–1.0 at 865 nm
Channels used in the retrieval	630 nm	650 and 850 nm	331 and 360 nm. Reported at 380 nm.	765 and 865 nm
Spatial resolution	Equal-area (110 km) <sup>2</sup>	AVHRR 4 km GAC (Global Area Coverage) data subsampled at 30 km	40×40 km	1 km
Cloud mask	A standard AVHRR clear/cloud classification (Stowe et al., 1999)	Modified ISCCP cloud detection scheme as described in Mishchenko et al. (1999)	Threshold of 360 nm reflectance and TOMS Aerosol Index information.	Uses threshold of the TOA reflectance at 865 nm (Rayleigh contribution corrected)

\* The differences between AVHRR-1 (2g) and AVHRR-1 (3g/QC) are (i) that an additional quality control is performed for AVHRR-1 (3g/QC) (Ignatov and Stowe, 2002b), (ii) two channels are used in the retrieval (630 and 830 nm), (iii) and another radiative transfer code is used to calculate the look-up-table (Vermote et al., 1997). The AVHRR-1 (3g/QC) is documented in Ignatov and Nalli (2002), Ignatov and Stowe (2002a), Ignatov et al. (2004a).

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**Table 1.** Continued.

Satellite instrument	VIRS	GOME/ATSR2 <sup>a</sup>	MODIS <sup>a</sup>	MISR <sup>a</sup>
Responsible scientists	A. Ignatov	T. Holzer-Popp and M. Schroedter	L. Remer	D. Diner, R. Kahn, J. Martonchik
References	Ignatov (2003); Ignatov and Stowe (2000); Ignatov et al. (2004b) <sup>3</sup>	Holzer-Popp et al. (2002a, b)	Tanré et al. (1997); Remer et al. (submitted, 2004) <sup>4</sup>	Martonchik et al. (1998, 2002) Kahn et al. (1998, 2001)
Aerosol microphysics	Same as for AVHRR-1	External mixing of 6 basic components from OPAC (Hess et al., 1998)	Bi-lognormal distribution created from choice of 4 fine and 5 coarse modes. All spherical. Refractive indices vary with mode and wavelength. $n_r=1.36-1.53$ . $n_i=0-0.005$ .	24 mixtures of up to 3 components, covering small, medium, and large, non-absorbing and partly absorbing particles, along with medium, and large, non-spherical mineral dust analogs
Channels used in the retrieval	630 and 1610 nm	415–675 nm (10 GOME bands)/658, 864 nm	533, 644, 855, 1243, 1632, and 2119 nm	672, 867 nm
Spatial resolution	Variable typically (10–20 km) <sup>2</sup>	80×40 km <sup>2</sup> / 1×1 km <sup>2</sup>	500 m reflectances used to create 10 km AOD product	17.6 km product, based on 16×16 regions aggregated from 1.1 km pixel data
Cloud mask	A CERES cloud identification developed by NASA/LaRC (Trepte et al., 1999)	1×1 km <sup>2</sup> combined VIS/IR thresholds APOLLO (Kriebel et al., 1989, 2003)	Spatial variability; IR tests: 1.38 micron test for cirrus. (Martins et al., 2002)	Multi-angle-based: Radiative camera-to-camera (RCCM) and Stereo-Derived cloud masks (SDCM), plus angular smoothness and spatial correlation tests at pixel resolution

<sup>a</sup> has also an aerosol retrieval over land

<sup>3</sup> Ignatov, A., Minnis, P., Loeb, N., Wielicki, B., Miller, W., Geier, E., Sun-Mack, S., and Trepte, Q.: Aerosol retrievals from TRMM/VIRS over open ocean, J. Appl. Meteorol., in preparation, 2004b.

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**Table 2.** Location of AERONET stations used in the comparison.

Name	Location	Altitude asl. (m)	Co-ordinates
Andros Island	Bahamas	0	N 24°41' W 77°47'
Anymon	Korea	47	N 36°31' E 126°19'
Arica	Chile	25	S 18°28' W 70°18'
Ascension Island	South Atlantic	30	S 07°58' W 14°24'
Azores	North Atlantic	50	N 38°31' W 28°37'
Bahrain	Persian Gulf	0	N 26°19' E 50°30'
Barbados	West Indies	0	N 13°09' W 59°30'
Bermuda	North Atlantic	10	N 32°22' W 64°41'
Capo Verde	North Atlantic	60	N 16°43' W 22°56'
Coconut Island	Pacific Ocean	0	N 21°25' W 157°47'
Dakar	Senegal	0	N 14°23' W 16°57'
Dongsha Island	South China Sea	5	N 20°41' E 116°04'
Dry Tortugas	Florida	0	N 24°36' W 82°47'
Goa	India	20	N 15°27' E 73°48'
Gotland	Sweden	10	N 57°55' E 18°56'
Helgoland	North Sea	33	N 54°10' E 07°53'
IMS METU ERDEMLI	Turkey	0	N 36°33' E 34°15'
Inhaca	Mozambique	73	S 26°02' E 32°54'
Kolimbari	Crete	0	N 35°31' E 23°46'
Kaashidhoo	Maldives	0	N 04°57' E 73°27'
La Paguera	West Indies	0	N 17°58' W 67°02'
Lanai	Hawaii	20	N 20°44' W 156°55'
Male	Maldives	2	N 04°11' E 73°31'
Nauru	Pacific Ocean	7	S 00°31' E 166°54'
NCU Taiwan	South China Sea	0	N 24°53' E 121°05'
Rame Head	England	0	N 50°21' W 04°08'
Roosevelt Roads	West Indies	10	N 18°11' W 65°35'
San Nicolas	California	133	N 33°15' W 119°29'
Shirahama	Japan	10	N 33°41' E 135°21'
Swakopmund	Namibia	250	S 22°39' E 14°33'
Tahiti	Pacific Ocean	98	S 17°34' W 149°36'
Venise	Italy	10	N 45°18' E 12°30'
Wallops	Virginia	10	N 37°56' W 75°28'

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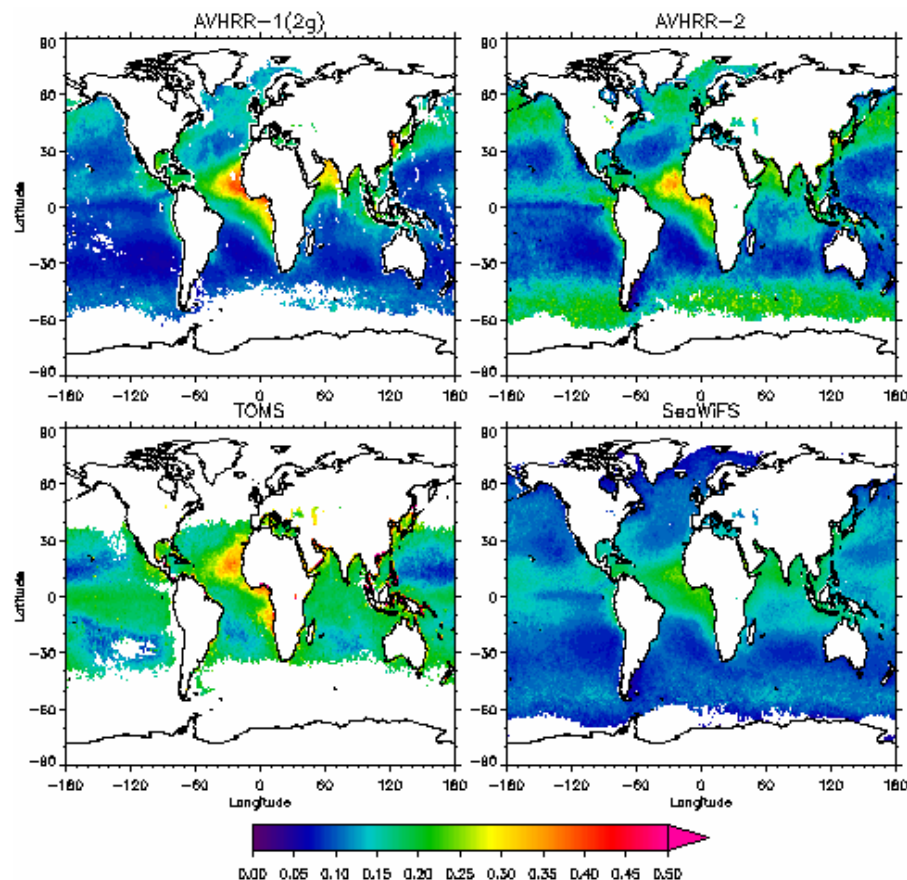
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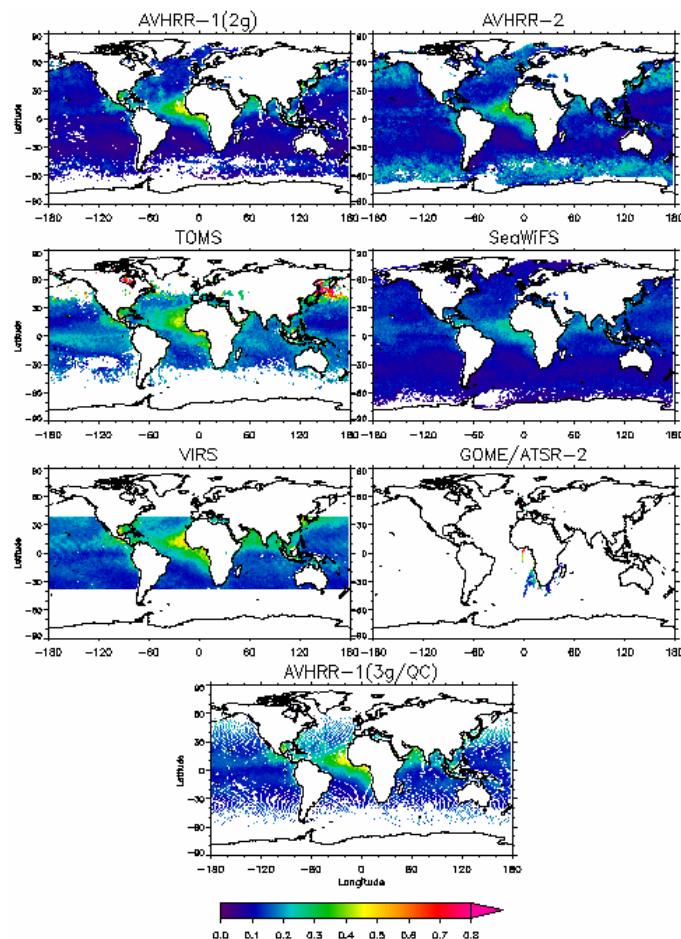


**Fig. 1.** (a) Averaged AOD (550 nm) over ocean for the period September 1997–December 2000. Values are given with minimum of data for ten months. Maximum AOD is 0.72, 1.00, 1.57, and 0.28, respectively, for AVHRR-1(2g), AVHRR-2, TOMS, and SeaWiFS.

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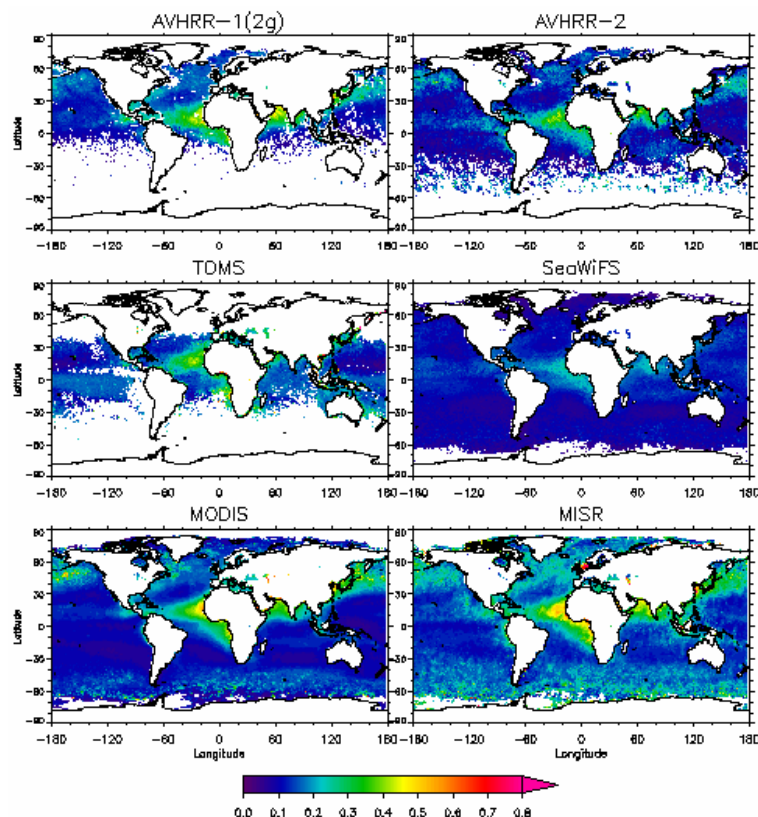


**Fig. 1. (b)** Averaged AOD (550 nm) over ocean for the period January to August 1998. Values are given with minimum of data for two months. Maximum AOD is 0.95 for AVHRR-1(2g), 1.00 for AVHRR-2, 2.53 for TOMS, 0.35 for SeaWiFS, 1.31 for VIRS, 2.73 for GOME/ATSR-2, and 0.64 for AVHRR-1 (3g/QC).

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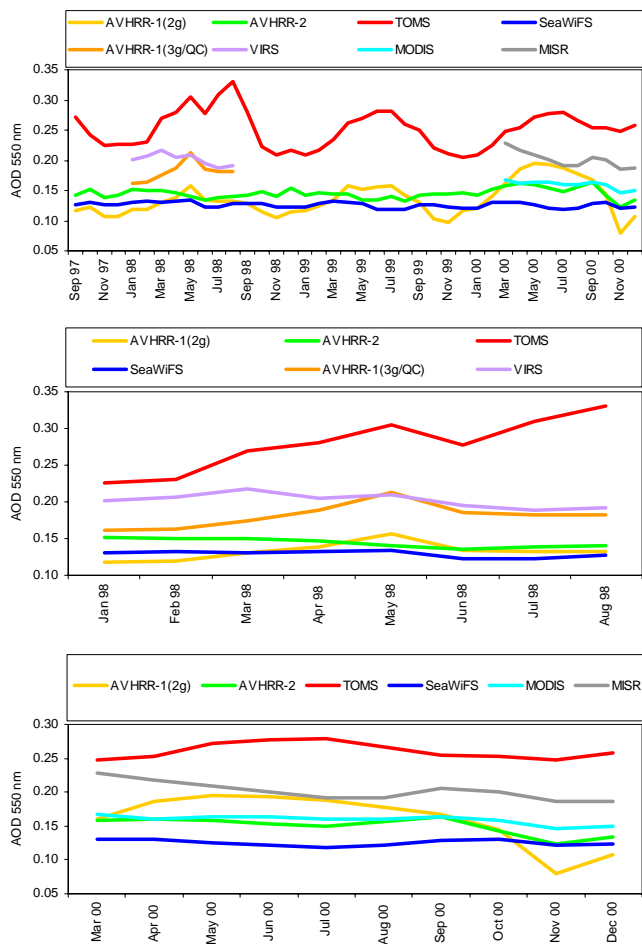
**Fig. 1. (c)** Averaged AOD (550 nm) over ocean for the period March to December 2000. Values are given with minimum of data for two months. Maximum AOD is 0.81 for AVHRR-1, 1.00 for AVHRR-2, 2.53 for TOMS, 0.31 for SeaWiFS, 2.27 for MODIS, and, 1.62 for MISR. (Note that orbit of NOAA-14 satellite (launched late 1994) whose AVHRR data are used in this study significantly drifted towards later afternoon by year 2000, causing loss of AVHRR-1 retrievals at low sun in the upper left panel.)

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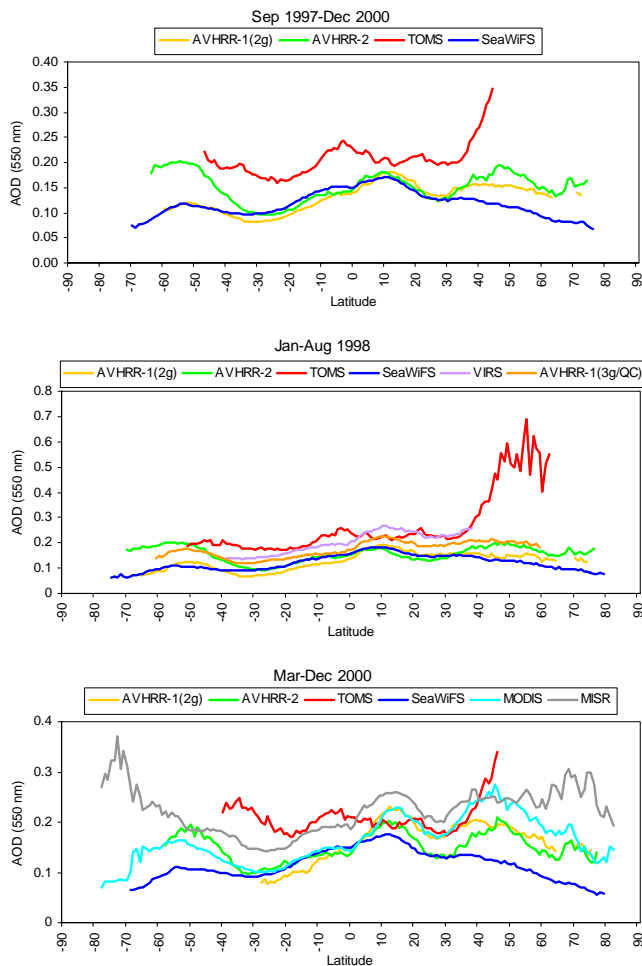


**Fig. 2.** (a) Global and monthly mean AOD (550 nm) over ocean from September 1997 to December 2000. (b) Global and monthly mean AOD (550 nm) over ocean for the eight-month period January to August 1998. (c) Global and monthly mean AOD (550 nm) over ocean for the ten-month period March to December 2000. 8233

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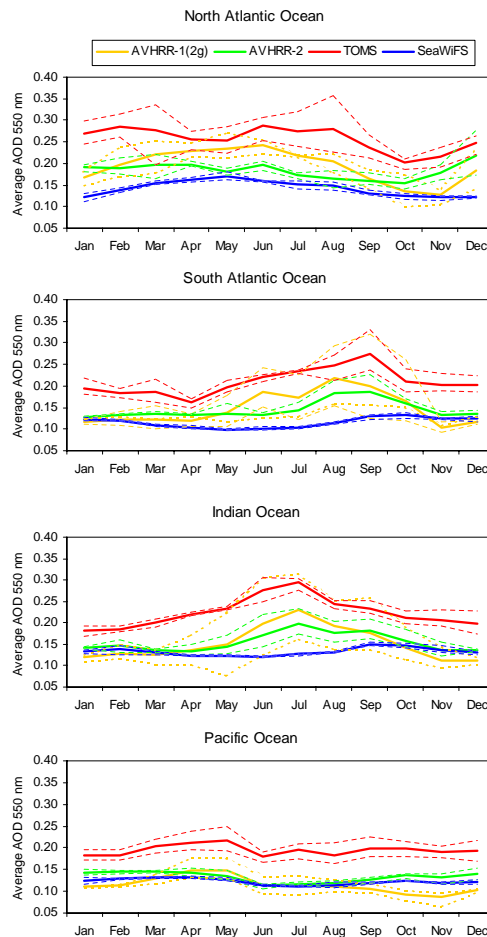


**Fig. 3.** Zonal mean AOD as a function of latitude, for the entire period of investigation, as well as for the two selected periods used in Fig. 1b and c.

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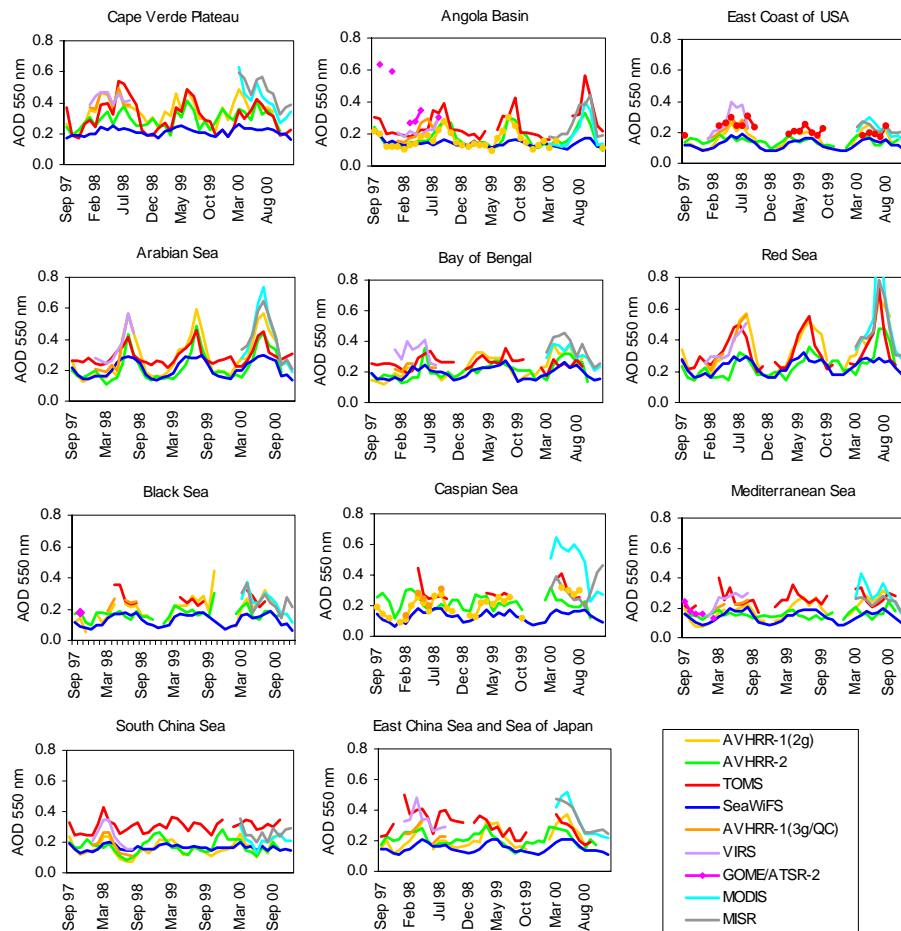


**Fig. 4.** Monthly mean AOD (550 nm) for four ocean regions (see definition in Fig. A2). The solid curves are averages for the entire period selected in this study. The dashed lines represents highest and lowest values among the years.

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**Fig. 5.** Monthly mean AOD (550 nm) over ocean for 11 near coastal regions (see definition in Fig. A2).

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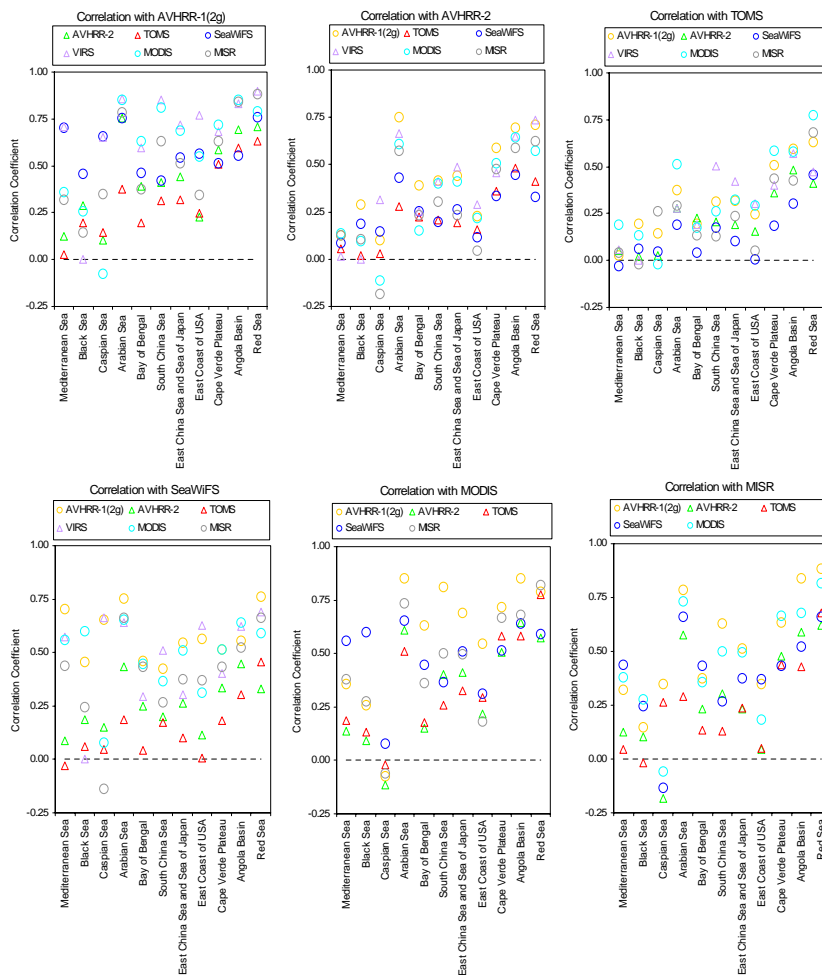
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**Fig. 6.** Temporal average of spatial correlations on grid square level between AVHRR-1(2g), AVHRR-2, VIRS, TOMS, SeaWiFS, MODIS and MISR for the 11 regions (see definition in Fig. A2).

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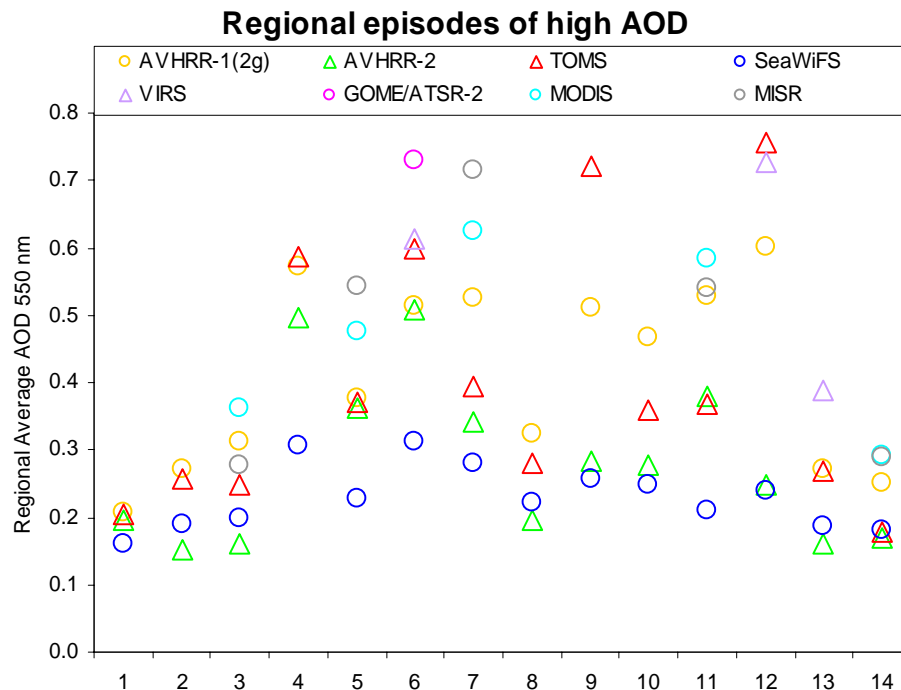
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**Fig. 7.** Comparison of AOD from satellite retrievals for episodes with high AOD and for periods of aerosol campaigns, 1) Europe Atlantic Coast, April 1999, 2) Mediterranean Sea, August 1999, 3) Mediterranean Sea, August 2000, 4) Cape Verde Plateau, June 1999, 5) Cape Verde Plateau, September 2000 (period of SHADE), 6) Angola Coast, August 1998, 7) Angola Coast, August 2000 (period of SAFARI2000), 8) Indian Coast, March 1999 (period of INDOEX), 9) Indonesia, October 1997, 10) China Sea, May 1999, 11) China Sea, March 2000, 12) Mexico, May 1998, 13) East Coast USA, May 1998, 14) East Coast USA, May 2000.

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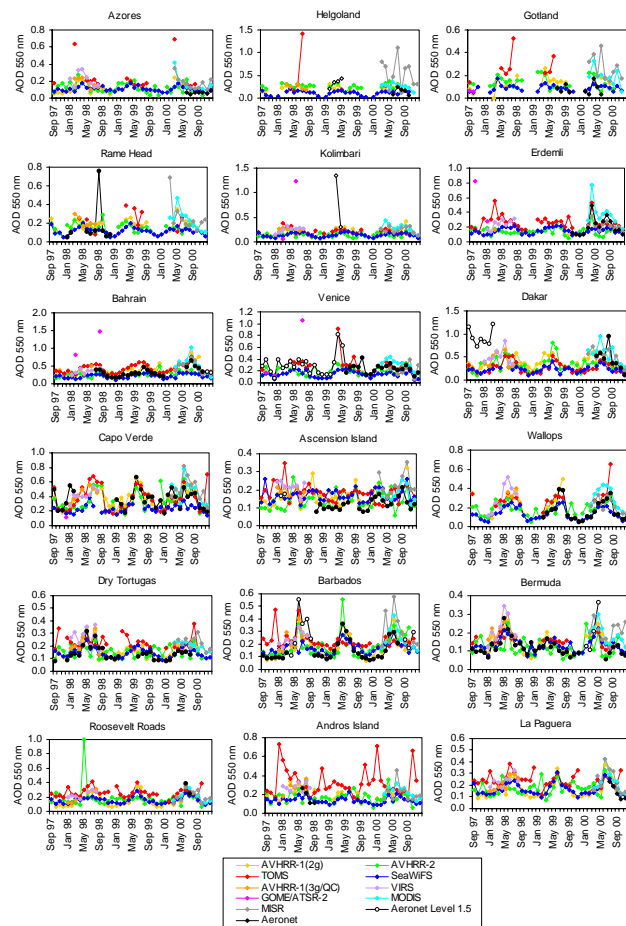
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**Fig. 8.** Monthly mean AOD from AERONET and nine satellite retrievals. The satellite data are for 550 nm, while the AERONET data are mean values of AOD at 440 and 670 nm (500 and 670 at some stations). Note the different scales for AOD at different stations.

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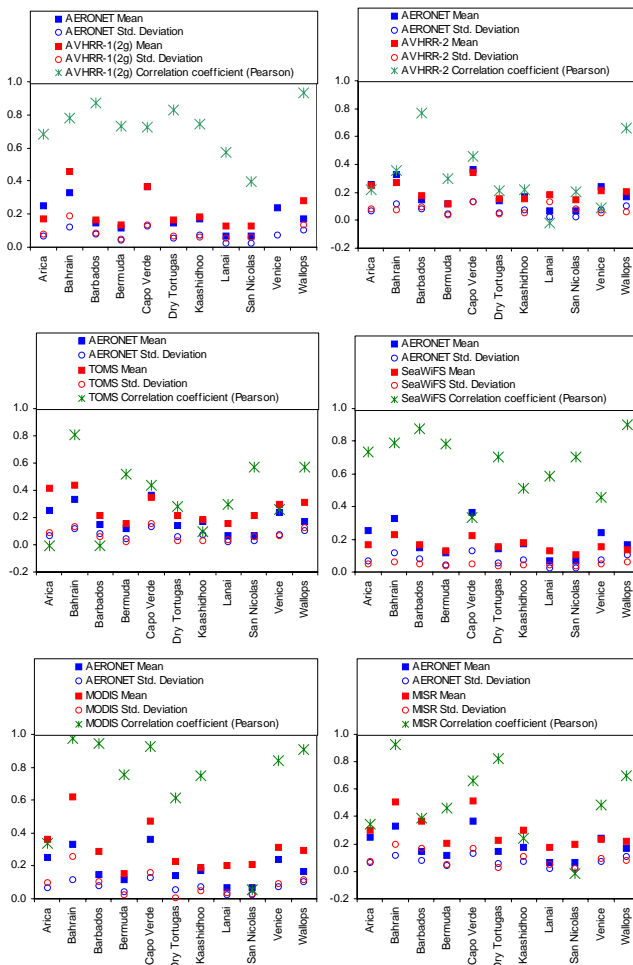
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**Fig. 9.** AERONET and satellite mean AOD and standard deviation and the correlation coefficient between the satellite retrieval and AERONET data. 11 stations and six satellite retrievals (one in each panel) are included.

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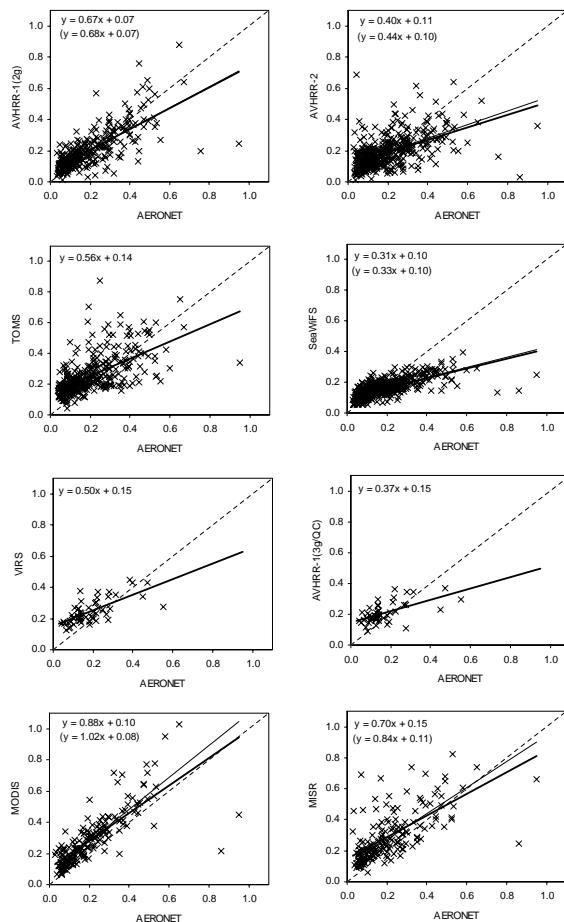
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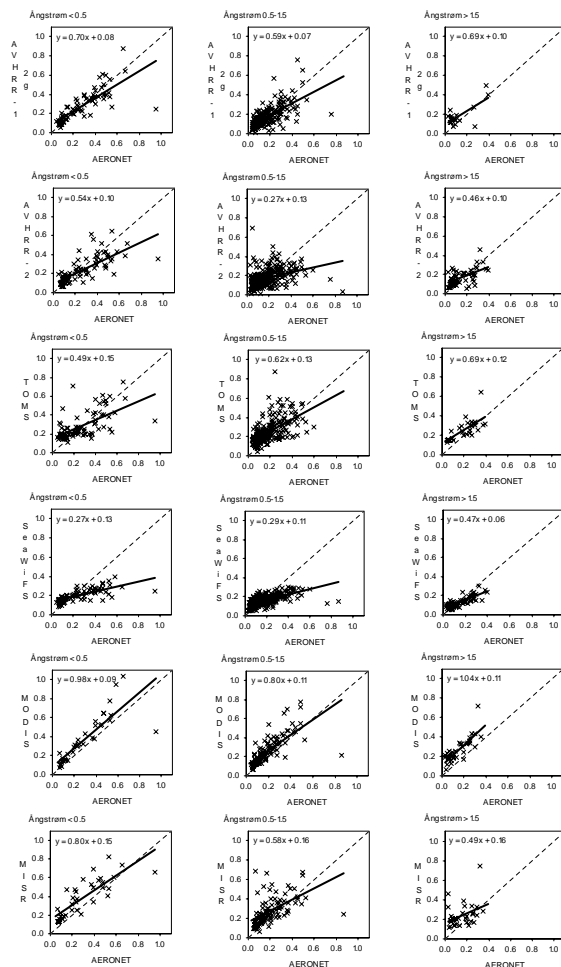


**Fig. 10.** Scatter plot of monthly mean AOD for AERONET data vs. eight satellite retrievals. Satellite data are given for 550 nm, AERONET data are mean of 440 (500 for some stations) and 670 nm. Results are also shown when values from Helgoland, Gotland (all data), and Taiwan October 2000 are removed from the analysis (thin lines, and equations in parenthesis).

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**Fig. 11.** Scatter plot of monthly mean AOD for AERONET data vs. satellite retrievals. The analysis is divided into three groups depending on the AERONET Ångström exponent.

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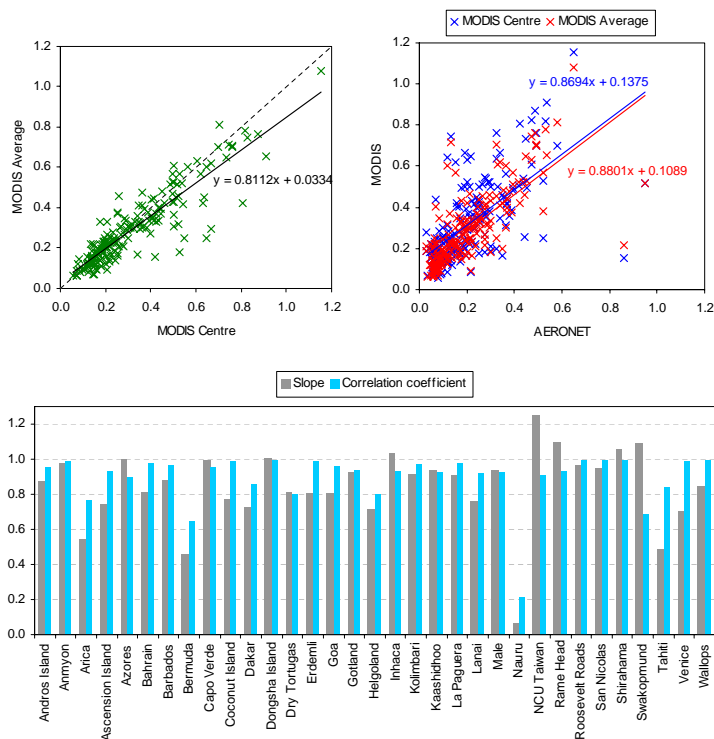
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**Fig. A1.** (a) Scatter plot of AOD in grid squares where AERONET stations are located (MODIS Centre) and average AOD for nine grid squares where the AERONET stations are located in the centre square (MODIS Average). (b) Scatter plot of AOD from AERONET and AOD from MODIS grid squares where the AERONET stations are located (MODIS Centre) and AOD average for nine grid squares where the AERONET stations are located in the centre grid square (MODIS Average). (c) Slope and correlation resulting from a regression analysis between AOD from MODIS grid squares where the AERONET stations are located (MODIS Centre) and AOD average for nine grid squares where the AERONET stations are located in the centre grid square (MODIS Average). Results are given for each of the AERONET stations.

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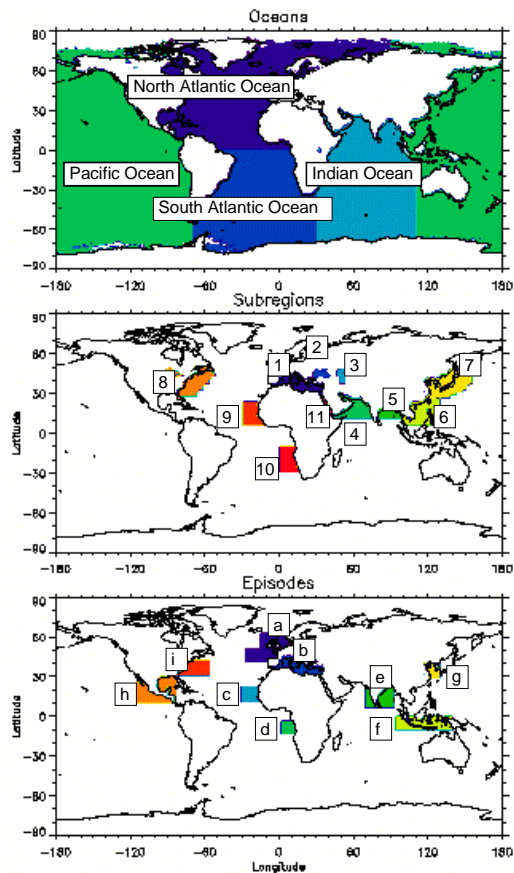
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- 2 Black Sea
- 3 Caspian Sea
- 4 Arabian Sea
- 5 Bay of Bengal
- 6 South China Sea
- 7 East China Sea and Sea of Japan
- 8 East Coast of USA
- 9 Cape Verde Plateau
- 10 Angola Basin
- 11 Red Sea

- a Europe Atlantic Coast
- b Mediterranean Sea
- c Cape Verde Plateau
- d Angola Coast
- e Indian Coast
- f Indonesia
- g China Sea
- h Mexico
- i East Coast USA

**Fig. A2.** Regions used in various analyses in this study (see text in Appendix B).

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