

54 **1. Introduction**

55 While clouds represent one of the largest modulators of Earth's radiation, with their impact
 56 dependent on a variety of cloud physical and radiative properties, they remain one of the
 57 more difficult components to represent in global climate models (Jiang et al. 2012).
 58 Passive satellite observational datasets such as those from MODIS (Moderate Resolution
 59 Imaging Spectroradiometer), AVHRR (Advanced Very High Resolution Radiometer),
 60 HIRS (High-spectral Infrared Sounder), and ISCCP (International Satellite Cloud
 61 Climatology Project) provide long-term, global cloud observations (Heidinger et al. 2013;
 62 King et al. 2013; King et al. 2003; Rossow 1991; Rossow; Schiffer 1999; Wylie; Menzel
 63 1999). However, assessing the uncertainties in the cloud radiative properties retrieved by
 64 these sensors has proved to be a complex and difficult task. Until recently, validation of
 65 these retrievals was limited to ground and aircraft inter-comparisons. But with the
 66 successful launch of CALIPSO (Cloud-Aerosol Lidar and Infrared Pathfinder Satellite
 67 Observations) and CloudSat in April 2006 as part of the NASA-led Afternoon
 68 Constellation (A-Train) (Stephens et al. 2002; Winker et al. 2010), researchers now have
 69 access to a near-continuous global record of vertically resolved observations of cloud and
 70 aerosol properties with nearly coincident observations from MODIS Aqua. Since launch,
 71 the CALIPSO lidar (the Cloud Aerosol Lidar with Orthogonal Polarization, or CALIOP)
 72 has proven to be a valuable tool for developing and evaluating passive cloud retrievals
 73 (Ackerman et al. 2008; Delanoë^{and}; Hogan 2010; Holz et al. 2008; Jin^{and} Nasiri 2013; Kahn et
 74 al. 2007). CALIOP can directly measure cloud-top height with sensitivities that ~~are an~~
 75 ^{significantly} ~~order-of-magnitude~~ greater than ~~the~~ passive retrievals, while the CALIOP depolarization

P = new paragraph

76 and attenuated backscatter measurements provide vertically resolved cloud phase
77 discrimination (Hu et al. 2009) for cloud layers up to a cumulative optical depth of about 3.

78 Ice Optical Thickness (IOT) has also proved to be one of the more challenging
79 properties to retrieve from space-based passive sensor measurement^s. In particular, it is
80 quite difficult to infer the microphysical and radiative properties of optically thin upper
81 tropospheric ice clouds (cirrus) from observations made by passive space-borne
82 instruments due to ~~the~~^{their} tenuous nature, extensive spatial scales, complex particle shapes,
83 and a wide range of particle sizes. There is a pressing need to conduct independent
84 validation to examine systematic biases between MODIS Collection 5 (C5) and CALIOP
85 Version 3 (V3) retrievals of tenuous IOT (< 3.0). To this end, we use a month of collocated
86 A-Train observation to compare the aforementioned retrieval products. A factor of two bias
87 is found between MODIS and CALIOP["] unconstrained["] retrievals (presented in Figure 1),
88 raising a major question regarding the utility of these data records to study ice cloud
89 radiative processes. Here, we seek to understand and resolve the CALIOP and MODIS IOT
90 biases.

91 Both MODIS and CALIOP IOT retrievals require a priori information concerning
92 the ice particle scattering properties that relate the measured reflectance (MODIS) or
93 attenuated backscatter (CALIOP) to the cloud's IOT and potentially the effective particle
94 size. MODIS ice cloud forward radiative calculations in the visible/near-infrared (VNIR)
95 depend directly on the ice particle phase function assumption, and to a first order on the
96 associated asymmetry parameter (g). For CALIOP, an assumed extinction-to-backscatter
97 ratio is required for ~~so-called~~^{so-called} unconstrained[®] retrievals where the algorithm is unable to
98 make reliable estimates of cirrus IOT by ~~measuring~~^{ratioing} the attenuated backscatter coefficients
99 in some clear air region immediately below cloud base (Young and Vaughan, 2009).

introduced
up
here

100 Because solar background signals greatly reduce the signal-to-noise ratio (SNR) of the
101 CALIOP daytime measurements, the vast majority of CALIOP daytime IOT estimates are
102 derived from unconstrained retrievals. ^PUncertainties in the ice scattering property
103 assumptions of either MODIS and/or CALIOP could account for the biases found in Fig. 1.
104 As will be discussed, an infrared (IR) cirrus IOT retrieval is relatively insensitive to ice
105 particle size and scattering details compared to MODIS and CALIOP VNIR measurements,
106 and thus provides an independent means to assess thin to moderately optically thick ~~cirrus~~ ^{ice-phase}
107 retrievals (IOT ~ 0–3). In addition, an IR retrieval provides radiative closure with solar
108 reflectance based ^{or} MODIS IOT retrievals in the sense that consistency in the two retrieved
109 IOTs also implies forward model consistency with the respective top-of-atmosphere (TOA)
110 VNIR and IR observations.

111 Using the NASA-funded SSEC Atmosphere Product Evaluation and Test Element
112 (PEATE), now re-named the Suomi-NPP Atmosphere Science Investigator Processing
113 System (SIPS), the sensitivity of MODIS retrievals to ice single ^{scattering} properties are
114 investigated by repeated analyses of collocated January 2010 CALIOP and MODIS
115 observations using a variety of ice crystal habits (Yang et al. 2012) and size distributions.
116 Based on comparisons against IR retrievals, the MODIS MYD06 Collection 6 (C6) ice
117 cloud optical property algorithm uses a single habit – severely roughened aggregated
118 columns (Yang et al. 2012) – instead of the size-dependent multi-habit model (Baum et al.
119 2005) used for C5. The MYD06 C6 results ~~also~~ compare well with a new CALIOP version
120 that uses a modified (larger) extinction-to-backscattering ratio for unconstrained IOT
121 retrievals.

* Slight confusion here regarding implied coordination of C5/V3 at L84. I see now you were merely stating what was "broken". But, just a word that I was confused, and that some text be added to better guide the reader toward your hypothesis.

122 **2. Ice Cloud Optical Thickness Retrieval Datasets**

123 An overview of the relevant retrieval methodologies is presented here with a focus on the
124 forward cloudy radiative transfer modeling assumptions and IR ~~cirrus~~^{IOT} optical thickness
125 retrievals developed specifically for this study.

126 **2.1 IR retrievals and radiative closure**

127 The MODIS channel suite includes a range of IR channels extending well into the CO₂
128 absorption region (13-15 μm). The calibration of the IR channels has been extensively
129 validated and proven to have high accuracy, with uncertainties less than 0.5 K across a
130 broad temperature range (Tobin et al. 2006). For ~~cirrus~~^{ice phase clouds}, the IR radiative transfer is
131 dominated by absorption, and thus is less complex than for the VNIR retrieval. In this
132 section, we discuss the IR radiative transfer methodology that is used both to retrieve the
133 IR IOT as well as evaluate the MODIS and CALIOP retrievals.

134 The goal of radiative closure study is to relate the differences in the CALIOP and
135 MODIS retrieved IOT to the measured TOA channel radiance or Brightness Temperature
136 (BT) in the MODIS 11 μm channel. To calculate the TOA cloudy radiances requires an
137 accurate radiative transfer model, knowledge of the cloud boundaries, and well-
138 characterized surface temperature/emissivity and atmospheric thermodynamic profiles.
139 LBLDIS (Turner et al. 2003), a cloudy radiative transfer model, is used for this analysis.
140 The model elegantly combines the clear sky Line By Line Radiative Transfer Model
141 (LBLRTM) (Clough; Moncet 1992) with the Discrete Ordinates Radiative Transfer
142 (DISORT) (Stamnes et al. 1988), a proven and accurate cloudy radiative transfer model.
143 The inputs required for LBLRTM are surface temperature and emissivity, vertically-
144 resolved temperature and water vapor profiles, and information regarding trace gas

145 concentrations such as CO₂ and O₃. For this analysis, ~~the~~ surface temperature and
146 thermodynamic profiles are extracted from the NOAA Global Data Assimilation System
147 (GDAS) files that provide profiles at 1° spatial resolution every 6 hours. For each
148 MODIS and CALIOP field of view (FOV), the closest (in both time and space) GDAS
149 profile is selected. A fixed CO₂ concentration of 380 ppm and a climatological O₃ profile
150 is used. Given these inputs, LBLRTM is run on the selected FOV filtered using the
151 collocated CALIOP V3 5km layer products (described in Section 3). The results of the
152 clear sky validation are discussed in Section 4

153 The cloud microphysics and thermodynamics are defined with a vertical
154 resolution of 500 meters within the cloud boundaries defined by the CALIOP layer
155 product. For example a cloud with a geometrical thickness of 1.5 km is divided into 3
156 layers, with each layer defined by an optical thickness, effective radius and ice scattering
157 model. For example, for a cloud with a total optical thickness of 1.5 each layer will have
158 an optical thickness of 0.5. Using this methodology, the vertical temperature profile is
159 accounted for in the radiative transfer. For daytime IR forward model calculations, the
160 effective radius from the MODIS optical property retrieval is used for all cloud layers.
161 For nighttime CALIOP comparisons, a fixed effective radius of 40 μm is used in the IR
162 calculations.

Therefore, you're assuming a 'block' cloud structure, rather than investigating potential signal/structure gradients?

are confused as to why 'night' is important here?

163 The last remaining variable needed to calculate the TOA IR radiance is IOT.
164 LBLDIS is run independently using either the MODIS or CALIOP retrieved IOT,
165 resulting in high spectral resolution TOA radiances with the only differences being the
166 assumed IOT (i.e., MODIS or CALIOP). The spectrally resolved radiances are then
167 integrated over the MODIS Aqua 11 μm channel (band 31) spectral response function

Are you constraining the CALIOP sample at all
to limit the effect of attenuation and overestimates
of ~~cloud~~ cloud base height?

168 resulting in a simulated TOA radiance that can be directly compared to the measured
169 MODIS 11 μm observations.

170 In addition to LBLDIS spectral calculations, TOA longwave fluxes are calculated
171 using the Rapid Radiative Transfer Model (RRTM) (Mlawer et al. 1997) that is also
172 based on DISORT and LBLRTM and utilizes a correlated- k method for gas absorption
173 along with broadband ice cloud parameterizations from Fu et al. (2000). Identical inputs
174 are used for RRTM and the LBLDIS TOA calculations with the only variable being IOT.
175 The TOA fluxes are subsequently used to quantify the impact of the IOT biases on the
176 global characterization of ice cloud radiative forcing.

177 IR observations provide the independent reference to understand differences
178 between MODIS and CALIOP IOT retrievals. While radiance closure provides valuable
179 information regarding TOA radiances and fluxes it does not provide a direct assessment
180 of the individual CALIOP and MODIS IOT biases. To convert observed IR TOA
181 radiance to IOT, two different retrieval approaches were used. First, we developed an IR
182 window IOT retrieval that uses the collocated MODIS and CALIOP observations. This
183 “reference” retrieval uses cloud boundary information from CALIOP coupled with the
184 LBLDIS forward model and then retrieves the IR IOT using the MODIS 11 μm window
185 channel observations that are coincident and collocated with CALIOP. A second method
186 uses the spectral emissivity retrieved from the MODIS CO₂ emissive cloud-top pressure
187 retrieval that is then related to the IOT and effective radius using a pre-computed lookup
188 table (Heidinger et al. 2015). This method has the advantage of being computationally
189 very efficient, not requiring the CALIOP cloud boundaries, and providing IOT for the
190 entire MODIS swath. Both IR retrieval methods are discussed in more detail in the
191 following sub-sections.

192 **2.1.1. Combined MODIS IR Window and CALIOP Retrievals**

193 A single channel IR window IOT retrieval was developed for this study using combined
194 CALIOP and MODIS observations and the LBLDIS forward radiative transfer modeling
195 discussed in the previous section. The method constrains the cloud boundaries using the
196 collocated CALIOP 5km layer products and uses surface and atmospheric temperatures
197 information from GDAS. TOA radiances are simulated using LBLDIS with IOT retrieved
198 by minimizing the measured MODIS channel 31 (11 μm) and calculated BT differences.
199 The retrieval assumes the cloud extinction is evenly distributed in the vertical throughout
200 the cloud. This simplification has the potential to bias the ~~IOT for FOV~~ ^{retrieval} where the IOT is
201 distributed non-uniformly in the vertical (Maestri; Holz 2009). The cloud geometric
202 thickness is thus limited to no greater than 4 km to reduce IOT biases that can be
203 introduced by non-homogeneous layers.

↳ Makes me worry even further about attenuation and resolving cloud base

204 **2.1.2. MODIS IR Spectral Emissivity Retrievals**

205 The MODIS C6 CO₂ slicing algorithm provides retrieved spectral emissivity for the 8.5,
206 11, and 12 μm channels (channels 29, 31, 32) that have sensitivity to both the IOT and
207 effective radius. As described in (Parol et al. 1991), β ratios can be approximated based
208 on these emissivities and are related to the asymmetry parameter (g), single-scattering
209 albedo (ω₀), and extinction efficiency (Q_e) as follows:

210

211 (1)
$$\beta_{\lambda_1\lambda_2} = \frac{Q_{e,\lambda_1}(1 - \omega_{0,\lambda_1}g_{\lambda_1})}{Q_{e,\lambda_2}(1 - \omega_{0,\lambda_2}g_{\lambda_2})}$$
 • ← Puckatur

212

213 Thus β is the ratio of the scaled absorption extinction in two spectral channels (λ_1 and λ_2).
214 The effective radius is first retrieved by matching simulated ice single-scattering
215 calculations of $g(r)$, $\omega_o(r)$, and $Q_e(r)$, each integrated over the appropriate MODIS
216 spectral response functions, to the retrieved MODIS β ratios. For this analysis the
217 scattering properties of severely roughened aggregated columns (Yang et al. 2012) are
218 used to be consistent with the MODIS C6 cloud optical property retrievals.

219 Using the effective radius to define $g(r)$, $\omega_o(r)$, and $Q_e(r)$, the extinction optical
220 thickness is then retrieved by relating the 11 μm emissivity to the extinction optical
221 thickness in the form ((Van de Hulst 1974))

222 (2)
$$\tau_{vis} = \frac{2}{Q_e} \left(\frac{\tau_{abs}}{(1-\omega_o g)} \right) ,$$

223 where τ_{abs} is the IR absorption optical thickness and τ_{vis} is the extinction optical
224 thickness at 532 nm. This derivation assumes that the ratio between the absorption and
225 extinction optical thickness is a factor of 2 in the IR. Based on ice cloud single-scattering
226 calculations (Yang et al. 2012), and assuming that the majority of ice clouds have an
227 effective radius greater than 10 μm , this assumption is expected to have introduced no
228 more than 10% uncertainty. (Heidinger et al. 2015) provides a more detailed discussion
229 of the retrieval methodology. This approach can be applied without the need for the
230 CALIOP cloud boundaries, and provides full swath IR IOT retrievals. We leverage this
231 capability to investigate the MODIS IOT retrieval biases as a function of view angle.

232 **2.2 CALIOP Ice Cloud Optical Thickness Retrievals**

233 CALIOP is a two-wavelength elastic backscatter lidar that measures attenuated
234 backscatter components polarized parallel and perpendicular to the transmitted laser light

235 at 532 nm and total attenuated backscatter at 1064 nm (Hunt et al. 2009). Once the
236 received signals have been background-subtracted and calibrated (Powell et al. 2009), a
237 tightly integrated suite of retrieval algorithms is used to detect layer boundaries (Vaughan
238 et al. 2009) and classify layers as either clouds or aerosols (Liu et al. 2009). Layers
239 classified as clouds are further classified according to thermodynamic phase as either ice
240 clouds or water clouds (Hu et al. 2009). Layer optical thickness (including IOT) is then
241 retrieved using one of two techniques: constrained or unconstrained retrievals (Young;
242 Vaughan 2009). Constrained retrievals are applied whenever the effective two-way
243 transmittance of a layer,

244

245 (3)
$$T_{eff}^2 = \exp(-2\eta\tau) = \exp\left(-2\eta \int_{layer\ top}^{layer\ base} \sigma_c(r) dr\right)$$
,

246

247

248 can be directly and reliably measured. In this expression, τ is the layer optical depth (IOT
249 for ice clouds), $\sigma_c(r)$ is the range-resolved cloud extinction coefficient, and η is a multiple
250 scattering correction factor whose value depends on the lidar sensing geometry and the
251 scattering characteristics of the particulates being measured. While T_{eff}^2 estimates can be
252 obtained from measurements of clear air, opaque water clouds, and ocean surfaces (see
253 (Josset et al. 2012; Yongxiang et al. 2007; Young 1995); respectively), the CALIOP V3
254 algorithm only implements the clear air technique, in which T_{eff}^2 can be obtained directly
255 from the ratio of the mean attenuated scattering ratios calculated in regions of clear air
256 located immediately above cloud top and below cloud base (Vaughan et al. 2005).
257 Retrieving IOT from measurements of T_{eff}^2 requires knowledge of the appropriate

order
of
reference →

Talk to
Mark about
Garner et al (2015)
and impact
on η

258 multiple scattering factor (Winker 2003). For CALIOP measurements of cirrus clouds,
259 (Josset et al. 2012) determined the mean multiple scattering factor to be 0.61 ± 0.15 . In
260 the CALIOP V3 algorithm, η is fixed at 0.6 for all cirrus clouds.

261 Constrained retrievals are the preferred method for retrieving IOT from CALIOP
262 measurements. However, because solar background light significantly degrades the
263 CALIOP SNR during daylight operations, V3 constrained retrievals occur almost
264 exclusively during nighttime observations, thus severely limiting direct comparisons with
265 MODIS IOT retrievals derived from VNIR solar reflectance. For the vast majority of
266 daytime observations, CALIOP IOT retrievals use an unconstrained technique that
267 requires *a priori* knowledge of the cirrus extinction-to-backscatter ratio (i.e., lidar ratio),

268 (4)
$$S_c = \frac{\sigma_c(r)}{\beta_c(r)}$$

269 where $\sigma_c(r)$ and $\beta_c(r)$ are, respectively, the cloud extinction and backscatter coefficients.
270 IOT is then obtained by solving the lidar equation using specified values of η and S_c
271 (Young; Vaughan 2009). Note that while the cloud extinction and backscatter coefficients
272 are explicitly range-dependent, their ratio is assumed to be range-invariant. Although S_c
273 for ^{ice layers} cirrus most likely varies depending on crystal habit and size distribution, the CALIOP
274 V3 unconstrained retrievals use a globally constant default value of $S_c = 25 \pm 10$ sr. This
275 value was determined prior to launch from the best information available from numerous
276 ground-based and airborne data sets (e.g., Holz 2002; Sassen 2001; Yorks et al. 2011).

277 Errors in lidar ratio selection for unconstrained retrievals generate corresponding
278 errors in the resultant estimates of IOT. In particular, an underestimate of S_c will result in
279 CALIOP underestimating IOT. The selection of the default CALIOP lidar ratio is thus
280 one of the potential major sources of bias in the CALIOP unconstrained retrievals that

281 can be investigated using IR observations from either MODIS or the CALIPSO IIR
282 (Imaging Infrared Radiometer) instrument (Garnier et al. 2015).

283

284 **2.4 MODIS Ice Cloud Optical Thickness Retrievals**

285 The MODIS imager provides measurements in 36 spectral channels, covering the Visible
286 Near Infrared (VNIR), Shortwave Infrared (SWIR), Midwave Infrared (MWIR), and
287 thermal IR portions of the spectrum. Spatial resolution is 250 m in two VNIR channels,
288 500 m in 5 VIS/SWIR channels, and 1 km in the remaining channels.

289 The MODIS cloud optical/microphysical property algorithm is used to generate a
290 single cloud product designated by the NASA Earth science data type (ESDT) names
291 MOD06 and MYD06 for Terra and Aqua MODIS, respectively (hereafter referred to as
292 MYD06 since the algorithms are essentially identical and this study is focused on
293 MODIS Aqua observations). For daytime measurements, the 1 km cloud retrieval
294 algorithm uses multiple spectral channels (primarily six VNIR, SWIR and MWIR
295 channels, as well as several thermal channels) to simultaneously retrieve cloud optical
296 thickness, effective radius (and derived water path) and thermodynamic phase for liquid
297 and ice phase clouds. In addition to the 1 km MODIS Level-1B calibrated radiance
298 product, the algorithm requires the following input: MODIS cloud mask (MYD35)
299 including 250 m mask information (Ackerman et al., 1998), the cloud-top pressure
300 portion of MYD06 (Ackerman et al. 2008; Holz et al. 2008), and a variety of ancillary
301 datasets. Heritage algorithm work is discussed in King et al. (2003), Nakajima; King
302 (1990), Platnick; Twomey (1994), Platnick et al. (2001); Platnick et al. 2003.

350 distinct layers in cases where the base of the upper layer is separated from the top of the
351 lower layer by as little as a single range bin (60m). For a passive retrieval such as from
352 MODIS, a 60 m vertical separation will have little impact on the retrieval results,
353 assuming both layers are ice. To improve the comparison yield and provide a more
354 representative distribution of single layer ice clouds for inter-comparing the passive
355 observations, CALIOP 5 km ice cloud layers with a vertical separation of 3 km or less are
356 merged to form single, vertically contiguous layers. The CALIOP extinction profile is
357 then integrated for each profile using the redefined layer boundaries, thus providing an
358 aggregated IOT. Ice clouds with total geometrical thickness greater than 4 km using this
359 single layer definition are excluded from the comparison.

Any idea how their decision then relates to use of block cloud profiles in IOT?

Hence the selection of 4 km depth at 202? yes!

360 The MODIS IOT retrievals are filtered using the C5 MODIS Quality Assurance
361 (QA) parameters and a horizontal heterogeneity threshold. MODIS IOT retrievals, (i.e.,
362 with the QA usefulness flag set to 1 and the QA confidence flag set to 3) are used in the
363 comparison. Using this filtering provides the highest quality MODIS retrievals and
364 removes all cloud edges from the comparison. To reduce uncertainties resulting from
365 spatial sampling differences between MODIS and CALIOP, the standard deviation of a
366 5x5 pixel box centered over the collocated pixel is computed. Only collocated pixels
367 where the MODIS IOT standard deviation is less than 0.5 are used; we find, however,
368 that the comparison results are relatively insensitive to this threshold.

C5?

369 Figure 1 reveals a systematic bias between the MODIS and CALIOP IOT's, with
370 MODIS approximately a factor of two larger than the CALIOP unconstrained retrievals.
371 An independent methodology is needed to assess this difference since both retrievals
372 depend on ice scattering property assumptions. As discussed in the methodology section,
373 the IR observations provide sensitivity to the IOT given well-constrained cloud

374 boundaries with uncertainties that are independent of the CALIOP and MODIS VNIR
375 retrievals. Spectrally resolved TOA radiances are calculated for the three different
376 retrieval methods – MODIS, CALIOP unconstrained (daytime measurements), and
377 CALIOP constrained (nighttime measurements) – using LBLRTM and LBLDIS. All
378 three calculations use identical cloud boundaries defined by the merged CALIOP 5 km
379 layer heights and the same thermodynamic profiles and ocean surface temperatures
380 (GDAS), with the only difference being the IOT used in the calculation. The spectrally-
381 resolved TOA radiances are then integrated over the MODIS channel 31 (11 μm) spectral
382 response function. To investigate the accuracy of the combined GDAS and TOA clear
383 sky LBLRTM calculations, simulated TOA 11 μm BT for clear sky ~~FOV~~ ^{collocations} identified
384 using both the MODIS and CALIOP cloud masks were compared to the measured
385 MODIS 11 μm channel BTs. The mean bias between the simulated and observed ^{clear-sky} BT is
386 less than 0.2 K, which is within the expected calibration uncertainty of MODIS (Tobin et
387 al. 2006).

~~Confused...
all these
clear MODIS
here?~~

388 Figure 2a presents the MODIS C5 and CALIOP V3 BT closure results. The figure
389 reveals a sobering finding which is that neither the MODIS C5 nor the CALIOP V3
390 unconstrained IOT retrievals provide radiative closure in the window IR. Furthermore,
391 the respective retrievals are biased in opposite directions. For MODIS C5, the calculated
392 TOA BT is colder than the measured BT with a mean bias of -8.7K, implying the MODIS
393 IOT is on average biased high. In contrast, the TOA BT calculated using the CALIOP V3
394 unconstrained IOT has a mean bias of +12.1 K, suggesting the CALIOP retrieval is
395 biased low. The CALIOP V3 constrained retrievals, which do not require an assumed
396 lidar ratio but only an estimate of the multiple scattering correction, demonstrate much
397 better agreement with a mean bias of +1.4 K.

* Is it this result, then, consistent with what you'd expect within the context of your available obs? MODIS IOT's are underconstrained w/ respect to diffuse cloud top elements, which you've directly resolved via CALIOP (= warm bias). CALIOP detection

398 To put the biases into a radiative context, the cloudy IR TOA fluxes are computed
399 for each collocated ^{ion}FOV using RRTM. The calculations use the CALIOP cloud
400 boundaries, the surface and atmospheric profiles from GDAS, and the MODIS retrieved
401 effective radius. For each collocated ^{ion}FOV, two RRTM calculations are computed with the
402 only difference being the IOT used (MODIS or CALIOP) with the results presented in
403 Figure 2b. The mean TOA flux difference between MODIS and CALIOP unconstrained
404 retrievals is $+23 \text{ W m}^{-2}$ with a standard deviation of 21 W m^{-2} . For the tenuous ^{ice phase layers} cirrus being
405 investigated, the sensitivity of the TOA flux to IOT is primarily driven by the thermal
406 contrast between the surface and the mean emitting temperature of the cloud. The very
407 large differences in the wings of the distribution in Fig. 1b occur primarily near the
408 tropics where the thermal contrast is greatest between the cloud and the surface. For this
409 region TOA differences as large as 50 W m^{-2} are found in Figure 2b.

of SCBH retrieval is questionable = warm bias?

REF < Corti and Peta 2009 is a good one.

410 5. IR Retrievals as a Reference Optical Thickness

411 Because the sensitivity of IR IOT retrievals to ice crystal habit selection is minimal, these
412 retrievals provide an independent means to evaluate the CALIOP and MODIS solar
413 reflectance retrievals. As discussed in Sect. 2, the main sources of uncertainty in the IR
414 IOT originate from characterizing the surface temperature and having an accurate
415 determination of the cloud emitting temperature. To reduce the surface temperature
416 uncertainty, the results of this section are restricted to non-polar (± 60 degrees) ocean-
417 only cases.

418 The comparisons with IR window IOT retrievals shown in Figure 3 reveal biases
419 in both the MODIS (a) and daytime CALIOP unconstrained (b) retrievals (high and low,
420 respectively) that are consistent with the radiative closure results presented in Figure 2.

421 The magnitude of the bias relative to the IR is approximately +40% for MODIS. For
422 CALIOP there is a non-linear dependence between the IOT and the negative bias relative
423 to the IR, with the bias increasing substantially for IR IOTs greater than unity; the
424 CALIOP results are discussed further in Section 5.2.

attenuation
of
CBH?

425 A limitation of the IR window IOT data set is that only a small subset of the
426 MODIS across track swath can be assessed due to the very close coordination between
427 the MODIS and CALIOP orbits. To investigate MODIS IOT scan angle dependencies, we
428 use the MODIS spectral IR IOT retrieval described in Sect. 2.1.2. Figure 4a shows the
429 MODIS C5 liquid (warm colors) and ice (cool colors) phase cloud optical thickness for
430 an example MODIS data granule (January 11 2010, 06:25 UTC). Fig. 4b presents the
431 histogram of the ratio between the MODIS IOT and the full swath IR IOT (described in
432 section 2.1.2) separated by viewing angle ranges as indicated by the colored lines
433 overlaid on the IOT image. A ratio of unity would suggest good agreement between the
434 spectral IR and VNIR IOT retrievals. However, as illustrated in the following section, for
435 the MODIS C5 retrievals (solid lines) the modes of the distributions vary with scan angle,
436 and the bias is seen to be an increasing as a function of scan angle. This is an important
437 result, as it demonstrates necessity that this scattering angle dependence can provide an
438 additional constraint on ice radiative model selection. In addition, because CALIPSO and
439 Aqua have similar orbits, only a small range of MODIS viewing angles are included in
440 the collocated inter-comparison, thus the possible strong dependence on viewing angle
441 implies the collocated analysis is representative only of the view angle ranges sampled.
442 Finally, given the lack of significant scattering in the IR, the scan dependent bias further
443 suggests the issue is with the MODIS C5 VNIR retrievals. This is investigated in the next
444 section.

and need for accurate cloud boundaries?

unclear.
Does it
show some
bias? If so,
does this
simply relate
to "effective"
IOT relative to
satellite FOV
that influences
your IR
inversion
retrievals?

445 **5.1 Ice Radiative Model Sensitivities in MODIS**

446 Though a primary focus of this investigation is on optimizing C6 ice models to
447 improve IOT intercomparisons, it is understood that ice model crystal habits also affect
448 the particle single scattering albedo retrieved using the SWIR and MWIR channels that
449 provide effective particle size information. Figure 5a and Figure 5b show the 2.13 μm and
450 3.7 μm channel co-albedo, respectively, as a function of Cloud Effective Radius (CER)
451 for four habit realizations; namely the C5 habit mixture (black line) and the three severely
452 roughened habits solid aggregate plates (green line), solid bullet rosettes (red line), and
453 aggregate columns (blue line). To the extent that CER retrievals of an asymptotically
454 thick cloud in the SWIR/MWIR are essentially a retrieval of co-albedo, the difference
455 between the aggregated column and C5 model co-albedo implies an effective radius
456 difference of +2 μm and -8 μm at the 2.1 μm and 3.7 μm wavelengths, respectively, for a
457 C5 effective radius of about 35 μm ; smaller C5 retrieved sizes would result in larger
458 differences.

459 Figure 6 shows the asymmetry parameter sensitivity to habit for the same four
460 habits shown in Figure 5. Evidently, the habit-sensitivity of the asymmetry parameter is
461 also strong in both the 2.1 μm and 3.7 μm MODIS channels. While the asymmetry
462 parameters of three severely roughened habits are not constant with effective size (though
463 at 2.1 μm the aggregate plates and aggregate columns are nearly constant), the C5 model
464 has much larger size sensitivity at both wavelengths. Aggregated columns, with smaller
465 asymmetry parameters relative to C5, will result in a larger retrieved CER estimates. This
466 is because the resulting increase in modeled SWIR reflectance for a given effective size
467 causes the measured reflectance to be associated with a more absorbing (i.e., larger)

* Why are sample sizes larger in Fig 7a for low IOT compared with Fig 3a?

468 particle. Therefore, the effect of both co-albedo and asymmetry parameter differences
469 between the severely aggregated column habit and the C5 model act to increase retrieved
470 effective radii at 2.1 μm , while at 3.7 μm some cancellation of effects can be expected.

471 The single habit radiative models shown in Figure 5 and Figure 6 are used to
472 build look-up tables that were integrated into the MODIS C6 cloud retrieval development
473 code. A month of data was processed for each habit. It was found that the habit that
474 provided the best consistency with the IR window retrievals (Sect. 2.1.1) is the severely
475 roughened aggregated column model. The IOT retrieval comparison with the IR window
476 retrievals using this model is shown in Figure 7a, where the MODIS reflectance-based
477 retrievals using the severely roughened aggregated column model are now clustered
478 around the 1-to-1 line. In addition, this aggregated column model was used to assess the
479 MODIS retrieval swath dependence previously shown in Figure 4b. The improvement of
480 the aggregated column model (dashed lines) relative to the C5 model (solid lines) is
481 significant. Both results led to the decision to use the severely roughened aggregated
482 column radiative model for the MODIS C6 cloud optical/microphysical property
483 retrievals.

484 Figure 8 shows an example of ice cloud retrievals for C5 and C6 for ^Ttyphoon
485 Fung-Wong. The typhoon was located south of Taiwan at the time of the MODIS Aqua
486 data granule acquisition on September 20, 2014 (0530 UTC). The C5 and C6 ice (cool
487 colors) and liquid (warm colors) cloud optical thickness retrievals are shown in the
488 middle and right panels, respectively. In addition to ice radiative model differences,
489 MYD06 C5 and C6 have different schemes for the cloud thermodynamic phase yielding
490 different ice and liquid phase pixel populations, though the optical thickness spatial
491 patterns are similar for regions having the same phase. Because of the different phase

492 assignments made by these two scheme, quantifying ice model retrieval sensitivities
493 requires the comparisons be restricted to only those pixels for which both algorithms
494 generate successful retrievals that identify identical cloud phases. With this pixel
495 filtering, the left panel of Figure 8b shows the normalized IOT distribution for the optical
496 thickness range of the plot. The C6 IOT mode is roughly 27% smaller than the C5 mode,
497 while the mean is decreased by about 15%, from 4.16 for C5 to 3.55 for C6. The 2.1 μm
498 ice cloud effective particle radius retrievals are shown in the right panel, with the C6
499 mode and mean both increasing by about 4 μm (+15%) for C6 relative to C5.

500 **5.2 MODIS C6 model selection methodology**

501 The MODIS IOT retrieval depends strongly on assumed ice scattering properties
502 that are needed to relate the measured reflectance to the retrieved IOT. The MODIS C5
503 retrieval used empirically derived habit and size distributions, with asymmetry parameters
504 ranging between 0.79 and 0.88, depending on the ice cloud effective radius (Baum et al.,
505 2005). By conducting an infrared closure analysis, we have shown that the C5
506 parameterization is not representative of the globally averaged ice scattering properties.
507 More recent investigations of the ice cloud asymmetry parameter suggest that most ice
508 clouds have values around 0.75 in the visible spectrum. Additionally, use of the C5 ice
509 cloud radiative model results in MODIS retrieval biases are strongly dependent on the
510 viewing angle, as demonstrated in Figure 4. These findings motivated the investigation of
511 new ice scattering models that have lower asymmetry parameters and weaker dependence
512 on ice effective radius. ^{REFS?} Since the MODIS C5 algorithms were finalized, new ice scattering
513 models that incorporate roughened ice crystal parameterizations have been developed
514 (Yang et al. 2012). Experimentation with these new models demonstrates that a modified
515 gamma distribution of severely roughened aggregated columns provides a significantly

516 lower visible asymmetry parameter (~ 0.75) that shows very little dependence on ice
517 effective radius. For testing purposes, the MODIS cloud retrieval algorithm team
518 implemented these new scattering properties in the MYD06 retrieval algorithm. The
519 updated algorithm was then run on the Atmospheric PEATE and the resulting data was
520 collocated with CALIOP measurements. Simulated TOA cloudy MODIS 11 μm brightness
521 temperatures (BT) ~~are~~ ^{were} then computed using the reprocessed MODIS IOT retrievals and are
522 compared to the MODIS measured BT, These new results presented in Figure 10b. ~~The~~
523 updated ice scattering models generate greatly improved IOT estimates that show very
524 close to a one-to-one correspondence with the independently derived IR IOT values
525 (Figure 7a). Additionally the view angle dependent bias is largely removed, as presented in
526 Figure 4c. Based on these results, the recently reprocessed MODIS C6 cloud
527 optical/microphysical property product (now in forward production) uses a modified
528 gamma distribution consisting of a single habit of severely roughened aggregated columns
529 for ice cloud retrievals. An additional benefit of the single habit is that it simplifies the
530 retrieval and increases the reproducibility of the scattering properties by the research
531 community.

There is no Figure 4c??
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532 Figure 10a presents the same filtered 2-D histogram comparing CALIOP and
533 MODIS as Figure 1 but using the ice radiative model modifications made for MODIS and
534 the updated lidar ratio (32 sr) for CALIOP. Figure 10b presents the IR radiative closure for
535 the update IOT retrievals for January 2010. Notice the large bias between the MODIS and
536 CALIOP un-constrained IOT is significantly reduced and the IR radiative closure shows
537 very good agreement for all three IOT retrievals. There is still a tendency for the MODIS
538 IOT to be larger than CALIOP in Figure 10a. The MODIS C6 IR closure in Figure 10b

?? where did this come from? } Comes on next page? out of order?
v. static η , how is this advance on Garnier et al (2015)?

539 also demonstrates this bias, with the tail of the distribution weighted to negative BT
540 differences suggesting the remaining bias is specific to MODIS.

541

542 **5.3 Ice Lidar Ratio Sensitivities in CALIOP**

543 As previously discussed, CALIOP uses one of two methods, (i.e., constrained and
544 unconstrained) to retrieve IOT. The constrained method requires high SNR in clear air
545 regions immediately above and below the cloud. This SNR requirement limits the
546 constrained retrieval primarily to nighttime ^{collocated} FOVs, because solar background light
547 severely degrades the clear air SNR during the daytime. This precludes direct comparison
548 of the constrained retrievals with the MODIS daytime optical property retrievals. The IR
549 retrieval, being day/night independent, allows for direct inter-comparisons between the
550 MODIS IR IOT retrievals and both the constrained and un-constrained CALIOP IOT
551 retrievals providing a means to evaluate the two retrieval methods against a consistent
552 reference.

553 Figure 3b presents the joint histogram between the unconstrained CALIOP and
554 the MODIS window IR IOT for January 2010 for single layer cirrus. The filtering criteria
555 are the same as in Figure 1, except both day and night observations are included. The
556 CALIOP layer optical thickness is filtered using the extinction quality control (QC) flags
557 provided as part of the L2 products. Only QC values of 0 (unconstrained solution, no
558 lidar ratio adjustment), 2 (unconstrained solution, lidar ratio decreased) and 4
559 (unconstrained solution, lidar ratio increased) were selected. Consistent with the findings
560 of (Garnier et al. 2015), Figure 3b shows CALIOP unconstrained IOT is significantly
561 low-biased with respect to the IR IOT, with a non-linear dependence as a function of

562 IOT. Figure 9 compares the CALIOP constrained retrievals (QC=1) to the MODIS IR
563 COT for the same filtering criteria. This comparison reveals a distinct difference between
564 the CALIOP constrained and unconstrained retrievals (Figure 3b), as the constrained
565 retrievals demonstrate a significantly smaller bias relative to the IR IOT. While the
566 CALIOP IOT retrieval requires estimates of the multiple scattering contributions for both
567 the constrained and unconstrained retrievals, the un-constrained method also requires an
568 assumed lidar ratio whereas the constrained retrieval does not. Because both retrievals
569 use an identical fixed multiple scattering factors, the difference between the constrained
570 and unconstrained retrievals relative to the IR can be attributed to the use of an assumed
571 lidar ratio in the unconstrained retrieval.

572 To investigate the sensitivity of the CALIOP IOT retrievals to the lidar ratio, a
573 month of CALIOP L2 products was processed (January 2010) with the default lidar ratio
574 increased to 32 sr. This revised value is the mean of all V3 constrained solutions of
575 randomly oriented ice clouds (3,091,952 cases) measured between 28 November 2007
576 (when CALIPSO permanently changed its pointing angle to 3° off nadir) and 30 June 2012.
577 It is important to note that the selection of this new default lidar ratio was based on on-
578 going quality assurance analyses conducted by the CALIOP algorithm team that were
579 wholly independent of the IR inter-comparisons. The modified CALIOP product was
580 ingested by the Atmospheric PEATE and collocated with both the MODIS C5 and C6
581 products and the MODIS IR retrievals. The modified CALIOP unconstrained retrievals
582 compared to the reference IR IOT is presented in Figure 7b. Compared to the standard V3
583 products (Figure 3b) the change in the lidar ratio significantly reduced the bias compared
584 to IR IOT, and the non-linear behavior at large IOT is almost completely removed. This
585 is because optical depth is a nonlinear function of lidar ratio, thus weakly scattering

586 layers show minimal changes in IOT while the changes in strongly scattering layers are
587 much more substantial. This result strongly suggests that the current V3 unconstrained
588 lidar ratio of 25 sr should be increased in future versions of the CALIOP data products.

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589 **6. Conclusions**

590 ~~The~~ MODIS Collection 5 (C5) ice optical thickness (IOT) retrievals are compared to the
591 ~~Version 3 (V3) CALIOP IOT for one month (January 2010) of collocated single layer ice~~
592 clouds. The comparison reveals a factor of two differences between the retrievals as
593 presented in Figure 1. Using IR ^{IOT retrievals based on MODIS 11 μ m observations} ~~observations from MODIS~~ as an independent means of
594 assessing the CALIOP and MODIS IOT clearly demonstrates that both retrievals have
595 significant biases, but in opposite directions: MODIS C5 systematically overestimates IOT
596 while CALIOP V3 systematically underestimates IOT.

597 The decision to use the single severely roughened aggregate column habit as the
598 MODIS C6 ice cloud radiative model was made solely to achieve closure with IR retrievals
599 in a global sense. Our use of this model for this purpose does not imply that it is a suitable
600 microphysical model for use in understanding ice particle physical processes, (e.g., size
601 distribution evolution, fall speed distribution, etc.) Furthermore, the IR comparisons were
602 done in conjunction with collocated CALIOP observations that allow for the filtering
603 of multi-layer ice phase clouds from the statistical study; ~~The~~ data set used here is clearly a
604 subset of actual scenes and so may not be reflective of the full distribution of ice clouds
605 observed by the sensors. Finally, it is recognized that using a fixed ice radiative model for
606 global retrievals is only meaningful in a climatological sense and may be expected to
607 breakdown in instantaneous and/or regional studies.

608 The severely roughened aggregated column model adopted for the MODIS C6 ice
609 cloud algorithm has a fixed aspect ratio with an asymmetry parameter of about 0.75 in the
610 visible for all effective sizes. This produces results that are quite consistent with those
611 generated using the Inhomogeneous Hexagonal Mono-crystal (IHM) model derived by
612 (C.-Labonnote et al. 2001) ~~with~~ ^{an} asymmetry parameter of about 0.77) that provided a good

613 match with observed POLDER view angle-dependent VNIR reflectance. Other studies
614 have also suggested that featureless (i.e., smooth) phase functions indicative of
615 roughened or highly asymmetric aggregated habits with relatively small asymmetry
616 parameters are needed to match aircraft and satellite observations e.g., (Baran et al. 2001;
617 C. -Labonnote et al. 2000; van Diedenhoven et al. 2013).

618 The Generalized Habit Model (GHM) (Baum et al. 2010) was also tested but did
619 not the same level of radiative closure with the IR IOT retrievals compared to the
620 severely roughened aggregated columns (comparison shown in Fig. 7a). While there was
621 an improvement with respect to the C5 ice model (comparison shown in Fig. 3a), the
622 GHM model resulted in IOT retrievals that were still significantly larger than the IR
623 because of larger asymmetry parameters in the visible relative to the severely roughened
624 aggregated column model (about 0.77 at an effective radius of 5 μm up to 0.82 at 60 μm).
625 (Cole et al. 2012) also tested the GHM as well as single habit models from (Yang et al.
626 2012) and (Yang et al. 2003) against POLDER polarized and total reflectance
627 observations across a range of scattering angles. Polarized angular observations agreed
628 well with a severely roughened version of the GHM. However, it was concluded that
629 there was no single habit/model that is best in all respects for the reflectance (derived
630 spherical albedo) consistency tests, though the severely roughened aggregated column
631 model was not included in the analysis. Similarly, (Baran; Labonnote 2007) also noted
632 that though the IHM model provided good consistency with POLDER directional
633 reflectance distributions, it was less successful in matching the angular distribution of
634 polarized reflectances. Due to vertical size stratification in ice clouds it is possible that
635 different models are needed to match polarized observations (weighted towards the
636 uppermost portion of the cloud-top) with total reflectance observations (weighted deeper

637 into the cloud), e.g., (Platnick 2000) and (Zhang et al. 2010). Given that MODIS
638 retrievals are based on total reflectance, it is expected that directional reflectance
639 consistency with POLDER is the more relevant metric. Further, the study of (Zhang et al.
640 2010) shows there is little difference between IOT retrieved from reflectance and IR
641 observations for the model case study considered. (Fauchez et al. 2014) demonstrated that
642 for 1km IR observations, sensitivities to 3-D effects are limited to horizontal
643 heterogeneity (plane-parallel approximation or PPA bias) and the effect of vertical
644 heterogeneity is small. Though more extensive heterogeneity studies are needed, these
645 studies do suggest the utility of using IR IOT retrievals to assess MODIS reflectance-
646 based ice radiative models. Finally, we note that recent comparisons have demonstrated
647 consistency between Aqua MODIS C6 IOT retrievals and those from AIRS Version 6
648 (Kahn 2015).

649 For CALIOP it is found that the bias relative to the IR for the V3 IOT retrievals
650 depends on the retrieval method used. While CALIOP can make direct measurements of
651 the effective two-way transmittance of the layer, the retrieved optical thickness depends
652 only on an estimate of the multiple scattering factor and the accuracy of the molecular
653 attenuated backscatter profile (calculated from a temperature and pressure profile using
654 Rayleigh scattering theory). However, daytime solar background noise limits the
655 applicability of this constrained retrieval technique to mostly nighttime observations, thus
656 prohibiting direct comparisons to the MODIS daytime optical retrievals. For the
657 constrained retrieval we find good agreement with the IR radiative closure (Figure 2) and
658 the IR IOT in Figure 9. However, the majority of the daytime CALIOP retrievals use the
659 unconstrained method that requires an *a priori* specification of the cloud extinction-to-
660 backscatter ratio. It is these unconstrained retrievals that are directly compared to the

661 MODIS C5 IOT in Figure 1 and to the IR in Figure 2 and Figure 3. The CALIOP V3
662 unconstrained IOT retrievals show a significant low bias relative to both the IR and the
663 constrained CALIOP retrievals. Since both CALIOP methods assume an identical multiple
664 scattering correction, this suggests that the default lidar ratio (25 sr) used in the V3
665 CALIOP unconstrained retrievals is too low. As part of this investigation the CALIOP
666 algorithm team processed a month of retrievals using a lidar ratio of 32 sr for the
667 unconstrained retrievals with results presented in Figure 7b. It is important to note that the
668 selection of a lidar ratio of 32 sr was not based on the IR inter-comparison studies, but
669 instead was derived from independent analyses of the nighttime constrained retrievals
670 conducted by the CALIOP algorithm team in order to improve the accuracy of the
671 CALIOP unconstrained retrievals and increase the consistency of IOTs reported by the
672 constrained and unconstrained retrievals.

673

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680 **Bibliography**

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