1	ACPD-15-24403-2015
2	The Latitudinal Structure of Recent Changes in the Boreal
3	Brewer-Dobson Circulation
4	by C. Shi et al.
5	
6	Text is formatted as follows: The referees' comments are written in bold.
7	Responses to the remarks are written in italic. Changes in the text of the manuscript
8	are marked with apostrophes (" ").
9	A copy of the revised manuscript including all the track changes is appended.
10	
11	Final response of the authors to the referees' remarks.
12	
13	We thank all reviewers for their helpful comments and we followed their suggestions
14	whenever possible. We feel that these changes have helped substantially to improve
15	the manuscript.
16	
17	
18	Following our detail responses to the remarks of the referees
19	
20	Anonymous Referee #1: P1-P10; Anonymous Reviewer #2: P11-P17
21	
22	
23	Anonymous Referee #1: P1-P10
24	Received and published: 13 October 2015
25	General comment:
26	1) At first: How the speed-up of the midlatitudinal downward branch results in
27	an increase of HCl is not further explained. The authors write with respect to
28	this: "Under the approximate conditions of no divergence and no horizontal

HCl gradient, the strength of the local downwelling dominates the 29 extratropical lower stratospheric HCl concentration by downward transport. 30 31 Therefore, the increase in the mid-latitudes and decrease in the Arctic of HCl after 2006/07 can support the local speedup in mid-latitude and slowdown in 32 the Arctic of the downwelling resulting from the narrowing equatorward of 33 the downwelling branch of BDC." It is not clear to me what the authors 34 really want to say with this statement. I might speculate that they wish to say 35 that high HCl amounts from the higher atmosphere are transported 36 downwards, but at the end I am left alone without any explanation. 37

38

Response: We mean that air parcels with high HCl concentration are transported
abnormally downward by the local speedup of downwelling branch in the
mid-latitudes. More explanation is in the section 3.4 of the revision L257-L280.

42

2) At second: The paper fully ignores all observational evidence that a positive 43 44 trend of age of air has been found for the Northern mid-latitudes during the last decade (roughly 2002 - 2012), and, from single balloon observations, even 45 for the last 30 years. This observational evidence comes not only from direct 46 47 observations of age tracers, but also models driven by re-analyses like ERA-I produce increasing age of air over these years. It would be necessary to 48 demonstrate how increasing age of air comes along with decreasing transport 49 times within an accelerated mid-latitudinal BDC branch - one possibility 50 would be significantly enhanced aging by mixing. The authors do not 51 comment at all on these aspects and make no attempt to discuss their findings 52 in the context of these earlier observations. 53

54

Response: The age of air derived from dynamics only indicates the total slowdown of the BDC but cannot show the structural change of the BDC. Hence, there is not a contradiction between our opinions and theirs. This is expressed in detail in the revision, in L257-L280.

3) At third: I am not able to follow the authors (and this might be my fault) with 60 61 their numerical analysis that finally allows them to arrive at their conclusions, and at statements like the following: "The local acceleration of the 62 mid-latitude downwelling results from the branch narrowing equatorward 63 which is related to weak planetary wave activity and cold polar vortex 64 enhancement." I am not aware of a mid-latitudinal downward branch of the 65 BDC that is separated from the downward branch in the polar regions, so 66 that I do not know how this narrowing can be caused, nor have I found any 67 demonstration in the paper that the planetary wave activity was weak(er) or 68 the cold polar vortex was enhanced (what exactly, was enhanced?). 69

70

71 **Response:** We showed the basic facts in Fig.3a and discuss in section 3.2, in 72 L169-L178.

73

74 Specific comments:

4) Abstract, p24404, L6/7: "... decreasing extratropical middle-lower
stratospheric HCl." I do not understand this statement. Weakening of the
BDC would result in longer transport times and, thus, an increase of HCl due
to longer exposure of CFCs to photolysis.

79

Response: HCl increasing in the mid-latitude stratosphere is only caused by the
strong downwelling because HCl was decreasing in the HCl source area when BDC is
weakening during 2001-2011. They are expressed in detail in the revision, L268-L280.

83

5) Abstract, p24404, L7-10: "However, the global ozone chemistry and related trace gas data records for the stratosphere data (GOZCARDS) show that the tropical lowermost stratospheric WV increased by 18%/decade during 2001–2011 and the boreal midlatitude lower stratospheric HCl rose 25%/decade after 2006." Is this your own analysis, or do you refer to

89	previous published material? In the latter case, a reference is required.
90	
91	Response: The specific trends are analyzed by ourselves and can be found in section
92	3.
93	
94	6) Abstract, p24404, L11/12:" a speedup of the mid-latitude downwelling." As
95	above, I do not understand this conclusion, since, to my understanding, a
96	slow-down of the BDC would indeed result in an increase of HCl in the
97	extratropics.
98	
99	Response: as in the revision, L268-L280.
100	
101	7) p24405, L5-8:"In addition, in December, January, February and March
102	(DJFM), a large proportion of WV transport into stratosphere over the
103	Tropical Western Pacific (TWP) is known as "stratospheric fountain" (Geller
104	et al., 2002; Bannister et al., 2004; Bonazzola et al., 2004). Thus, the
105	variability of the BDC in the boreal winter affects the annual tropical lower
106	stratospheric WV (Dessler et al., 2014)." The BDC is understood as a zonally
107	averaged phenomenon. In this sense it is not correct to assign longitudinally
108	restricted processes like the transport over the Western Pacific to the BDC.
109	
110	Response: We deleted these redundant sentences in L51-L45.
111	
112	8) p24405, L19/20: "However, the contrary HCl trends after 2006 in the
113	mid-latitudes and the Arctic in boreal middle stratosphere (Fig. 2d and f)
114	cannot be explained simply by the slowdown of the downwelling argued by
115	Mahieu et al. (2014)." This is a bad organisation of the paper. Figs. 2d and f
116	are referred to in the introduction that already show result of the analysis

here - the introduction is not the correct place for this. Beyond this, Figures
need to be referred to in the order of their appearance.

Response: This sentence is a comment on Malieu et al. So we deleted our figures here
121 and revised to "However, the opposing HCl trends in the mid-latitudes and the Arctic
122 in boreal middle stratosphere in Fig. 4a of Mahieu et al. (2014) cannot be explained
123 simply by the slowdown of the downwelling.", in L66-L69.

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9) p24405, L26 and throughout the paper: "air age" - the term commonly used
is "age of air" (in full: stratospheric mean age of air, abbreviated: AoA)
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Response: We revised all of them.

10) p24406, L1/2: "Additionally, the HCl concentrations can be different in the air parcels with the same age but different transport pathways (Waugh et al., 2007)." This fact is accounted for in the concept of mean age of air by the age spectrum - indeed depends the mean age of air on the shape of the age spectrum since it is the first moment of the spectrum. Two air parcels with the same mean age but representing different transport pathways would have different age spectra that - only by chance - could have the same first momentum. Not impossible, but not very realistic. My impression is that the authors are not familiar with the concept of the age spectrum?

Response: These tedious sentences in context were removed in L78-L80.

11) p24408, L14-16: "During 2001–2011, the temperature near the tropical
tropopause (10S–10N) (Fig. 1b) rises with a trend of 2.21 K/decade based on
the GOZCARDSMERRA data." Is this consistent with other data sets, eg.
SPARC temperature trend analysis? It seems to be a tremendous and
unrealistically high amount!

Response: A similar trend in ERA-interim is 2.13K/decade at 70 hPa. We did not find



12) p24408, L19-22: "In addition, the decline of BDC during 2001-2011 was 152 confirmed by air age increases in the boreal main stratosphere from 153 154 Lagrangian transport models and observations (Ploeger et al., 2015; Mahieu et al., 2014; Stiller et al., 2012)." The increase and decrease of AoA shows a 155 rather complicated pattern in these publications; I do not agree with this 156 general statement made here; this argument mixes observations of AoA 157 trends in the Northern mid-stratosphere with trends of uplift velocities in the 158 tropical lower stratosphere. 159

160

162

13) p24408, L24: "There is not only a shift of the BDC trend ..." What kind of
shift? In what domain? There was no mentioning of any kind of shift before
in the paper.

166

Response: The trend shift near 2000 in Fig. 1a is from increase to decrease.

168

14) p24409, L2: "... the flow would be accelerated if the flow pipe narrows"
Where has it been shown that the "flow pipe" narrows? What is meant with
the "flow pipe"? The tropical pipe?

¹⁶¹ *Response:* The redundant sentence was removed in L160-L163.

173 **Response:** as in L174-L178, " If we regard the BDC as a flow pipe, the flow would be 174 accelerated when the downwelling branch narrows following the principle of mass 175 conservation. Fig.3a supports the pattern of the equatorward narrowing of the 176 downwelling branch, a speedup in the mid-latitudes and a slowdown in high latitudes 177 under the background of an enhanced polar vortex."

178

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179 15) p22409, L6-27: This entire paragraph is a mystery to me. I cannot follow
180 where these structural changes in the BDC are deduced from.
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181

182 *Response:* We revised section 3.2, as in L169-L205.

183

16) p22409, L13-15: "In other words, the downwelling branch narrowed
equatorward and sank faster in the mid-latitudes (downward black thin
hollow arrow in Fig. 1d)." I cannot follow at all. My understanding of the
BDC is that BOTH in midlatitudes and polar latitudes downwelling is the
predominant direction of the BDC; what do the authors want to say with
their statement? That downwelling occurs only in the polar vortex regions?

190

Response: as in comment 15. "According to the TEM momentum equation for the middle atmosphere, the weak wave activity may enhance the zonal flow (Fig. 2b) and weaken the BDC. Climatologically, the BDC downwelling branch is affected by zonal flow and is located throughout the stratosphere south of the jet axis and partly in the lower stratosphere under the jet core on the north of the jet (shading in Fig.3a). ", as in L169-L205.

197

198

17) p22409, L15/16: "The regression of EPFp from Fig. 1a with temperature of
DJFM in 2001–2011 (Fig. 1e) can ..." What has been plotted over what, and
which kind of regression analysis has been done? I cannot follow what has
been done here.

Response: as in L194-L198, "Using single variable linear regression of temperature on the mean EPF (Fig.3b), the positive temperature anomalies (near 0.5K) in the middle latitudes and negative anomalies (near -2 K) in the Arctic can also support the latitudinal structural changes of the downwelling branch during the period, the equatorward narrowing of the downwelling branch, a local speedup of downwelling in the mid-latitudes and a slowdown in the Arctic. "

210

18) p22410, L1-4: "The local acceleration of the mid-latitude downwelling results
from the branch narrowing equatorward which is related to weak planetary wave
activity and cold polar vortex enhancement." Again, I do not understand what is
meant with "branch narrowing equatorward". Where has it been demonstrated that
the planetary wave activity is weak, and where has the cold polar vortex
enhancement been shown?

217

Response: We revised section 3.2, as in L170-L210.

219

19) p22410, L15-16: "Fueglistaler (2012) analyzed tropical stratospheric WV
from HALOE and also found the trend after 2000." Since HALOE
terminated its operation in 2005, the trend derived by Fueglistaler is certainly
for another time period than the trend in this paper?

224

Response: as in L228-L232. They was revised to "Randel et al. (2006) and Dhomse
et al. (2008) indicated the opposite variation of WV during several years including
2000, a distinct mutational year (Fueglistaler, 2012). Here we extend the data series
and discuss WV during 2001-2011 after the mutation."

229

230 20) p22410/11, section 3.4: I understand this section as follows: since HCl vmr
231 increases with altitude, a faster downward transport in mid-latitudes brings
232 down higher HCl amounts from above, while in the Arctic, where the

233 downward transport is not accelerated according to the authors, the decrease of HCl continues in line with the decrease of CFCs. This interpretation is 234 purely speculative. The authors fully ignore in their argumentation that there 235 is independent observational evidence of increasing age of air in the Northern 236 mid-latitudes. This is in clear contradiction to their result of accelerated 237 downward transport in Northern midlatitudes. Even without this 238 239 contradictory observational evidence, the claims of the authors would need to be manifested by some thorough assessments on the amount of HCl increase 240 due to pure transport versus HCl decrease due to air becoming younger and 241 less photolyzed. Besides this, faster downward transport does not necessarily 242 mean that air from higher up in the atmosphere (where HCl vmr is larger) is 243 transported down. 244 245 Response: We revised this section, please refer to L257-L280. 246 247 248 21) p24411, L18/19: "The trends of boreal BDC tended to decrease and create 249 latitudinal structural changes." Where has this been shown in this paper? 250 251 Isn't this merely a summary of previous publications? By the way: do the authors really mean that the trend of the boreal BDC has changed, or the 252 **BDC itself?** 253 254 Response: Fig.1a shows the negative trend of the BDC after 2000 and and Fig. 3a 255 256 shows the structural changes. Refer to L169-L185. 257 22) p24411, L23: "...weaker planetary wave activity and the stronger polar vortex 258 after 2000." Where exactly has this been shown in the paper? 259 260 Response: We revised this section. Please refer to L169-L216. 261 262

263	23) p24411, L26/27: "The increasing HCl in the midlatitudes is caused by the
264	local speedup of downwelling after 2006/07." This interpretation is in clear
265	contradiction to the Mahieu et al. paper. The exact mechanism for this
266	increase is not explained at any place in this paper. The reader might assume
267	that the increased downward motion might bring higher HCl vmr from
268	higher altitudes down to the middle stratosphere. However, the competing
269	process of reduced photolysis during shorter transport times due to
270	accelerated BDC has not been assessed. Without this assessment of competing
271	processes the claim of the authors remain purely hypothetical.
272	
273	Response: We revised section 3.4, as in L257-L280.
274	
275	24) p24412, last para of the paper: The last para consists of a few random
276	citations about possible future evolution of the BDC that does not help at all
277	for the argumentation in this paper here.
278	
279	Response: The para was deleted in L300-L317.
280	
281	Technical comments:
282	25) Fig. 1: Is there any reason to squeeze all these panels within one single figure?
283	
284	Response: We reorganized the figures.
285	
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294 Anonymous Referee #2: P11-P17

295 Received and published: 26 October 2015

296

297 General Comments:

In this study authors try to explain anonymous changes stratospheric HCl and H2O over last few years. Authors use observation-based GOZCARDS data and ERA-interim dynamical fields to argue that there are significant changes in the stratospheric circulation (or Brewer-Dobson (BDC) circulation). They argue that weakening on tropical upwelling caused by TTL warming (hence increasing trend in the stratospheric H2O) but enhanced downwelling caused increase in mid-latitude stratospheric HCl.

305 Some aspects of this study are quite good but overall authors fail to explain their results more clear way. There is no limit on number of figures or words. So 306 authors should include detailed analysis of each aspect of their results. Most of 307 308 the sections/sentences seem to combination of different ideas and there is no clear flow in their arguments. Presenting all the analysis in just couple of figures also 309 does not help. Some references are cited just for sake of it and some are cited 310 311 only after reading the abstract. Results presented in Mahieu et al (2014), clearly show increase in mid-latitude age of air throughout the NH stratosphere. Their 312 analysis does not show any differences between the shallow branch and the deep 313 branch circulation. Can you please explain why your results differ as you are 314 also using ERA-interim data. Authors also forgot to include careful analysis of 315 H2O trends presented in Hegglin et al., 2014. Overall I think this manuscript 316 needs major revision before it is accepted to ACP. 317

318

319 Minor Comments

320 Page 24404

3211. Line 5-6 Confusing sentence: "Climatologically, a symmetric weakening BDC322indicates increasing tropical lower stratospheric WV and decreasing

extratropical middle-lower stratospheric HCl"

324

Response: These were revised to "Therefore, a symmetric weakening BDC would increase the tropical lower stratospheric WV and decrease the extratropical middle-lower stratospheric HCl", in L19-L21.

328

329 2. Line 14-15: enhancing polar vortex? Do you mean strength of polar vortex? 330 But when? early winter/mid-winter or late winter?

331

Response: as in L28-L32, "Results present that the enhancing polar vortex and weaken ing planetary wave activity leads to a downwelling branch narrowing equatorward and a local speedup of 24% at 20 hPa in the mid-latitudes." was revised to "Results present the accelerated winter mean Arctic circumpolar westerlies and the weakened planetary wave activity, which led to a equatorward narrowing of the downwelling branch and a local speedup of 24% at 20 hPa in the mid-latitudes.". The winter mean is in all DJF.

339

340 3. Line 16: What is regressive temperature increase? Where is your regression 341 model. What terms are included?

342

Response: "there are regressive temperature increase of 1.5K near the tropical tropopause and that of 0.5K in the midlatitude middle stratosphere, " was revised to "Using a single variable linear regression of temperature on the mean EPF, the temperature anomalies show a positive 1.5K near the tropical tropopause and a positive 0.5K in the mid-latitude middle stratosphere", in L32-L36

348

4. Line 23: aren't you contradicting yourself with Solomon et al., 2010?

350

Response: We agree with solomon, as in L42-L44, "In fact, WV decreased sharply

near 2000 and then increased mildly in the tropical lower stratosphere. Solomon et al.

353	(2010)".
354	
355	Page 24405
356	5. line 21: Do you know why?
357	
358	Response: The redundant sentence was deleted.
359	
360	Page 24406:
361	6. line 2 and 3: Confusing link between Waugh et al., (2007) and deep BDC.
362	Please revise
363	
364	Response: The tedious sentences in context were removed in L78-L81.
365	
366	7. line 16-21: Can you please comment on various biases in GOZCARDS data.
367	For careful and detailed analysis WV trend using satellite data, please see
368	Hegglin et al., 2014 (Nature Geoscience)
369	
370	Response: "The errors of the average GOZCARDS HCl and WV are less than 15% in
371	most of the stratosphere (Froidevaux et al., 2015). Hegglin et al. (2014) merged other
372	satellite data sets with the help of a chemistry-climate model and they determined
373	there is a link between the trends of stratospheric WV and BDC during 1986-2010.",
374	in L102-L106.
375	
376	8. line 24-26: Using monthly mean data to calculate EPF and TEM is pointless.
377	You should use daily fields and then calculate monthly mean values
378	
379	Response: EPF trend from daily fields as the following figure is similar to the EPF
380	trend from the monthly mean interim data as Fig.1a.



2001-2011 in this paper and did not discuss the period of 1993-1999 in detail. Page 24410 12. line 10-15: WV trends? Do you think Randel et al., 2006 or Dhomse et al., 2008 are incorrect to show increase BDC caused decrease in stratospheric water vapour after 2001? Response: in L228-L232, "Randel et al. (2006) and Dhomse et al. (2008) indicated the opposite variation of WV during several years including 2000, a distinct mutational year (Fueglistaler, 2012). Here we extend the data series and discuss WV during 2001-2011 after the mutation. " **Figures** 13. Figure 1a: For which years linear lines are fitted? **Response:** 1993-1999 and 2001-2011 14. Figure 1b: Why 25 month smoothing? **Response:** to remove seasonal and QBO signal. 15. Why do you want to have all the analysis in two figures. Please separate EP flux analysis as a new Figure **Response:** We reorganized the figures. 16. Can you please explain if the EP flux analysis shown in Figure 1 are from NCEP or ERA-interim. Do you use daily fluxes or just use monthly mean fields. Also I think better to use anomalies, not the absolute values.

436	Response: "Time series of 45°N-75°N mean vertical EPF from ERA-Interim in DJF at
437	50hPa. The monthly mean fluxes have the similar trend to that from daily fields.
438	Moreover, the anomalies and original values have the same trends.
439	
440	17. Plot U winds anomalies and EP flux anomalies separately.
441	
442	Response: U and EPF are usually plotted on one figure for analyzing their
443	coordinated variations.
444	
445	18. Figure 1f and 1h- 3 and 5 year running means. I assume you are using earlier
446	years from ERA-interim for pre-2001 time period but how are you truncation
447	last 5 years. Those end points must be skewing the time series
448	
449	Response: We only used the years during 2001-2011 and the running mean value
450	series does not occupy the start and the end points.
451	
452	19. Figure 2a: Can you please comment on quality of H2O data from
453	GOZCARDS. As it is combination of SAGE/HALOE and MLS. But as soon
454	as you add/remove satellite data, GOZCARDS seems to show strange
455	behaviour. I think authors should carefully read GOZCARDS related
456	document and should comment on those biases?
457	
458	Response: This data was widely used. Moreover, we can get similar figures to some
459	other papers, analyzing the data we used. As in L102-L106, "The errors of the
460	average GOZCARDS HCl and WV are less than 15% in most of the stratosphere
461	(Froidevaux et al., 2015). "
462	
463	20. What residuals are almost similar to absolute values. Your regression model
464	seems to have some problem. 21. Why do you use only solar term?

Response: The residuals were added by constant terms. Thus they were similar to absolute values. The other terms even the solar term cannot pass a high statistical significance test. So, in the revision, we gave up the regression of HCl and only used the original values.

470

1 The Latitudinal Structure of Recent Changes in the

2 Boreal Brewer-Dobson Circulation

3 4

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13

14 Abstract

The uUpwelling branch of the Brewer-Dobson Circulation (BDC) controls the 15 tropical lower stratospheric water vapor (WV) through dynamic cooling near the 16 tropopause. The dDownwelling branch of BDC dominates the extratropical 17 middle-lower stratospheric Hydrogen Chloride (HCl) by dynamic transport. 18 ClimatologicallyTherefore, a symmetric weakening BDC would increase indicates 19 increasing the tropical lower stratospheric WV and decreasing decrease the 20 extratropical middle-lower stratospheric HCl. However, the global ozone chemistry 21 and related trace gas data records for the stratosphere data (GOZCARDS) show that 22 the tropical lowermost stratospheric WV increased by 18% decade⁻¹ during 23 2001-2011 and the boreal mid-latitude lower stratospheric HCl rose 25% decade⁻¹ 24 after 2006. We interpret this as resulting from a slowdown of the tropical upwelling 25 and a speedup of the mid-latitude downwelling. This interpretation is supported by 26 27 composite analysis of the Eliasen-Palm Flux (EPF), zonal wind and regression of 28 temperature on the EPF from the ERA-Interim data. Results present the accelerated winter mean Arctic circumpolar westerlies and the weakened planetary wave activity, 29 which led to that the enhancing polar vortex and weakening planetary wave activity 30

leads to a equatorward narrowing of the downwelling branch narrowing equatorward and a local speedup of 24% at 20 hPa in the mid-latitudes. Moreover, Using a single variable linear regression of temperature on the mean EPF, the temperature anomalies show a positive 1.5K near the tropical tropopause and a positive 0.5K there are regressive temperature increase of 1.5K near the tropical tropopause and that of 0.5K in the mid-latitude middle stratosphere, which is also indicativees of the tropical upwelling slowdown and the mid-latitude downwelling speedup during 2001-2011.

38

39

1 Introduction

It has recently been discovered that the stratospheric trace gases, such as Water 40 Vapor (WV) and Hydrogen Chloride (HCl), tend to have anomalous trends since the 41 beginning of the 21st century (Mahieu et al., 2014; Fueglistaler, 2012). In fact, WV 42 decreased sharply near 2000 and then increased mildly in the tropical lower 43 stratosphere. Solomon et al. (2010) suggested that decrease of the lower stratospheric 44 WV near 2000 leads to a 25% reduction of the global warming rate near 2000, which 45 indicates that WV plays an important role in the balance of the global radiation. Mote 46 et al. (1995) detected the 'tape recorder' effect on tropical WV above the tropopause 47 and pointed out that the WV variability is dominated by dehydration based on the 48 tropical Cold-Point Tropopause Temperature (CPTT). The tropical CPTT is 49 connected with dynamic cooling of upwelling branch of Brewer-Dobson Circulation 50 (BDC). In addition, in December, January, February and March (DJFM), a large 51 proportion of WV transport into stratosphere over the Tropical Western Pacific (TWP) 52 is known as 'stratospheric fountain' (Geller et al., 2002; Bannister et al., 2004; 53 Bonazzola et al., 2004). Thus, the variability of the BDC in the boreal winter affects 54 the annual tropical lower stratospheric WV (Dessler et al., 2014). 55

The principal source for Stratospheric HCl is the decomposition of chlorofluorocarbons (CFCs) in the upper stratosphere under intense ultraviolet radiation (UV). HCl is relatively long-lived in the stratosphere (Mohanakumar, 2008; Andrews et al., 1987) and is transported downward by the downwelling branch of <u>the</u> deep BDC. Due to the Montreal Protocol, the total atmospheric chlorine <u>fromof</u> CFCs

decreased since 1992/1993 and stratospheric HCl declined after 1997 as shown from 61 ground-based Fourier-transform infrared (FTIR) spectrometers and various satellite 62 63 observations (Kohlhepp et al., 2011; Jones et al., 2011; Froidevaux et al., 2015). 64 When CFCs emission tends to be stable, HCl concentration in the upper stratosphere is dominated by the photochemical exposure which is also related to deep-upwelling 65 of BDC (Waugh et al., 2007). However, the contrary opposing HCl trends after 2006 66 in the mid-latitudes and the Arctic in boreal middle stratosphere (Fig.2d and Fig.2f) in 67 Fig. 4a of Mahieu et al (2014) cannot be explained simply by the slowdown of the 68 downwelling argued by Mahieu et al (2014). There might be the opposite opposing 69 trends of the local downwelling in the mid-latitudes and the Arctic. 70

Therefore, the combination of the variations of WV and HCl can be used to fully 71 understand the change of the deep BDC which is abbreviated as BDC in the below. 72 Recent studies found that the BDC has weakened since the beginning of the 21st 73 century in the Northern Hemisphere (NH) according to air agethe age of air (Ploeger 74 et al., 2015; Stiller et al., 2012). However, the mid-latitude air age in the middle and 75 lower stratosphere is not only dominated by the deep BDC but also by the shallow 76 BDC. Furthermore, the deep BDC and the shallow BDC can vary independently 77 (Garny et al., 2011; Gerber, 2012; Stiller et al., 2012). Additionally, the HCl 78 concentrations can be different in the air parcels with the same age but different 79 transport pathways (Waugh et al., 2007). Therefore, the deep BDC has direct and 80 efficient influence on the stratospheric HCl and WV. Hence, the mismatcheding trends 81 of the WV in the tropics and HCl in the extratropics in the recent observations allow 82 83 one to deduce an abnormal latitudinal structural change in the BDC. This kind of 84 latitudinal structure changes have has not been discussed before.

Hence, we focused on two points in this study: 1) the latitudinal structure of the recent boreal BDC; 2) the connections between the latitudinal structure of the BDC changes and the <u>mismatching-mismatched</u> trends of WV and HCl. The data set is described in section 2. Section 3.1 and 3.2 present the trends of the boreal BDC and the latitudinal structural changes of the BDC. Section 3.3 shows the relationship between BDC and WV. Section 3.4 discusses the connection between BDC and HCl. 91 The conclusion is provided in section 4.

92 93 2 Data and methodology 94 **2.1 Data** Monthly zonal average WV, HCl and temperature data are from Global OZone 95 Chemistry And Related trace gas Data records for the Stratosphere (GOZCARDS 96 Version 1.1, 1991 to 2012) satellite data set. WV data is derived from Level 2 satellite 97 products by HALOE (1991-2005) V19, UARS MLS (1991-1993) V5, ACE-FTS 98 (2004-2012) V2.2 and Aura MLS (2004-2012) V3.3. HCl data is derived from Level 2 99 products by HALOE (1991-2005) V19, ACE-FTS (2004-2012) V2.2 and Aura MLS 100 (2004-2012) V3.3. Temperature data is derived from MERRA (1979-2012) V5.2.0. 101 102 Details can be obtained from http://gozcards.jpl.nasa.gov. The errors of the average GOZCARDS HCl and WV are less than 15% in most of the stratosphere (Froidevaux 103 et al., 2015). Hegglin et al. (2014) merged other satellite data sets with the help of a 104 chemistry-climate model and they determined there is a link between the trends of 105 106 stratospheric WV and BDC during 1986-2010.

European Centre for Medium-Range Weather Forecast (ECMWF) ERA-Interim reanalysis monthly mean data (Dee et al., 2011) with $1.5^{\circ} \times 1.5^{\circ}$ grid was used to calculate <u>the</u> Eliasen-Palm Flux (EPF) and Transformed Eulerian Mean (TEM) velocity.

111

112 **2.2 Methodology**

113 EPF of planetary waves (PWs, wave number 1-3) is expressed in equation (1) and 114 TEM velocity (residual meridional circulation, approximative approximating theo 115 BDC) is expressed in equation (2) (Andrews et al., 1987). In vector Figures of EPF 116 based on ERA-interim data, items are is normalized by 3.14×6378 km horizontally and 117 by 1000 hPa vertically when in a vector figure. The units is are m²s⁻² 118 (http://www.esrl.noaa.gov/psd/data/epflux/).

$$\begin{cases} F_{(\varphi)} = -r_0 \cos \varphi u' v' & (1.a) \\ F_{(p)} = fr_0 \cos \varphi \frac{\overline{v' \theta'}}{\overline{\theta}_p} & (1.b) \end{cases}$$

$$\left(\overline{v}^{*} = \overline{v} - \frac{\partial}{\partial p} \left(\frac{\overline{v'\theta'}}{\overline{\theta_{p}}}\right)$$

$$\overline{\omega}^{*} = \overline{\omega} + \frac{\partial}{r_{0}\cos\varphi\partial\varphi} \left(\frac{\overline{v'\theta'}\cos\varphi}{\overline{\theta_{p}}}\right)$$
(2.a)
(2.b)

121 Where o Overbars and primes denote zonal means and departures, subscripts without 122 parentheses denote partial differentiation, p is the pressure, φ is the latitude, f123 is the Coriolis parameter, r_0 is the radius of the earth, θ is the potential temperature, 124 u and v are zonal and meridional velocity, and φ is vertical velocity.

Deep BDC is driven by the breaking of PWs (EPF convergence) in the winter 125 126 extratropical stratosphere-in-winter as known as the 'downward-control principle' (Haynes et al., 1991; Holton et al., 1995). Due to tThe vertical EPF expressed as-in 127 equation 1b (EPFp, the eddy heat flux) indicates the upward propagation of PWs 128 mainly in the extratropical stratosphere., This implies the EPFp in the lower 129 stratospherice region 45°N-75°N could be a proxy for the BDC (Newman et al., 2001; 130 Li and Thompson, 2013). Generally, EPFp is negative. Thus, the absolute value of 131 EPFp at 50hPa in 45°N-75°N can-will represent the intensity of the BDC. 132

133 December-February (DJF) mean is used for boreal winter EPF, $\overline{\omega}^*$ and other 134 dynamic variables. However, temperature is averaged from December to March 135 (DJFM), because the influence of PW activities in DJF on stratospheric temperature 136 will continue till March (Newman et al., 2001).

After <u>using a 25-month</u> moving average, the high-frequency signals are eliminated, <u>theand</u> long-term (more than two years) interannual variability is left. <u>Using single variable linear regression of temperature on the mid-latitude</u> <u>lower-stratospheric EPF, the correlation between temperature and planetary wave</u> <u>activity is discussed. Linear regression is used to calculate the correlation between</u> <u>temperature and EPFp.</u>

143

144 **3 Results**

145 **3.1 Trends of the boreal BDC**

The BDC proxy (EPFp at 50 hPa over 45°N-75°N in Fig.1a) in DJF from 146 ERA-Interim shows that BDC in boreal winter tended to increase in 1993-1999 period. 147 However, the BDC proxy tends to weaken with a trend of 2.21 $\times 10^5$ pa \cdot m⁻² \cdot s⁻² \cdot 148 decade⁻¹ (33% decade⁻¹) after 2000, which passes statistical significance test at 149 the 90% confidence level. The change of upwelling in the BDC can affect the tropical 150 CPTT by dynamic heating (i.e. less upwelling in the tropics implies less cooling and 151 greater heating). Thus, the tropical lowermost stratospheric temperature can be used to 152 153 test the long-term trend of the BDC. During 2001-2011, the temperature near the tropical tropopause (10°S-10°N) (Fig.1b) rises with a trend of 2.21 K decade⁻¹ based 154 on the GOZCARDS-MERRA data. A similar trend with 2.13 K decade⁻¹ is also 155 detected at 70 hPa in ERA-interim (figure not shown). The trend exceeds the 156 significance test at the 99% confidence level. The tropical tropopause warming is 157 consistent with the declining decline of the BDC observed in Fig.1a. Apparently, the 158 result is similar to the ozone analysis in the tropical lower stratosphere (Aschmann et 159 al., 2014) that where the BDC tends to decrease after 2000. In addition, the decline of 160 BDC during 2001-2011 was confirmed by air age increases in the boreal main 161 stratosphere from Lagrangian transport models and observations (Ploeger et al., 2015; 162 Mahieu et al., 2014; Stiller et al., 2012). 163

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- 165

3.2 The latitudinal structural changes of the BDC

166 There is not only a trend shift (near 2000 in Fig. 1a) of the BDC-trend but also latitudinal structural changes of the BDC after 2000. BDC (red thick hollow arrow in 167 Fig.1c2a and 3a) is a wave-driving net Lagrangian transport of mass through the 168 middle atmosphere (Andrews and McIntyre, 1976). According to the TEM 169 momentum equation for the middle atmosphere, the weak wave activity may enhance 170 the zonal flow (Fig. 2b) and weaken the BDC. Climatologically, the BDC 171 downwelling branch is affected by zonal flow and is located throughout the 172 stratosphere south of the jet axis and partly in the lower stratosphere under the jet core 173 on the north side of the jet (shading in Fig.3a). If we regard the BDC as a flow pipethe 174 principle of mass conservation, the flow would be accelerated if where the flow pipe 175

narrows following the principle of mass conservation. Fig.3a supports the pattern of 176 the equatorward narrowing of the downwelling branch, a speedup in the mid-latitudes 177 and a slowdown in high latitudes under the background of an enhanced polar vortex. 178 Fig.1f 3c shows an acceleration of the downwelling (also as shown by 'downward 179 black thin hollow arrow' in Fig.1d and 1e) on the south side of axis of the polar 180 vertex axis (35°N-about 65°N in Fig.1c2a) corresponding to the polar vortex 181 enhancement (in Fig.1d) at 20 hPa. The 2007-2011 average TEM vertical speed of 182 4.02×10^{-4} pa · s⁻¹ increased by 24-% over 2001-2005 average TEM vertical speed 183 of 3.25×10^{-4} pa \cdot s⁻¹. In contrast, the downwelling was decelerated at 20 hPa on 184 the north side of the polar vertex axis (Fig.3d). (in Fig.1f). Comparing to the averaged 185 EPF (BDC proxy) and zonal wind in winter during the period of 2001-2011(Fig.1c), 186 the differences of EPF and zonal wind between 2007-2011 and 2001-2005 show that 187 the decline in the upward propagation of PWs could result in a negative anomaly of 188 BDC (red thick hollow arrows in Fig.1d) accompanied by a polar vortex enhancement. 189 It is suggested that the more powerful circumpolar westerly induced by weaker PWs 190 191 would block the downwelling branch of the BDC outside the polar region. In other words, the downwelling branch narrowed equatorward and sank faster in the 192 mid-latitudes (the equatorward side of the polar vortex). (downward black thin hollow 193 arrow in Fig.1d). The regression of EPFp from Fig.1a with temperature of DJFM in 194 2001-2011 (Fig.1e) Using single variable linear regression of temperature on the 195 mean EPF (Fig.3b), the positive temperature anomalies (near 0.5K) in the middle 196 latitudes and negative anomalies (near -2 K) in the Arctic can also support the 197 latitudinal structural changes of the downwelling branch during the period, .- A 198 negative center (near -2 K) in the Arctic implies an upwelling anomaly (upward black 199 thin hollow arrow) which is validated by the weakening downwelling and shifting to 200 upwelling (in Fig.1g) on the north side of axis of polar vertex (65°N-85°N) at 20 hPa. 201 The anomalies of the temperature and TEM vertical velocity in the Arctic area 202 indicate the equatorward narrowing of the downwelling branch was narrowing 203 equatorward. The narrowing leads to , a local speedup of downwelling in the 204 mid-latitudes and a slowdown in the Arctic. - (Fig.1f and the downward black thin 205

hollow arrow in Fig.1e) which is also verified by a positive center (near 0.5K) in the
middle stratosphere. The positive center (near 1.5K in Fig.1e-3b accounts for 62% of
the total temperature increase in Fig.1b) is located in the lowermost stratosphere over
tropics, from which the reduction of the upwelling of the BDC can be inferred
(downward red thick hollow arrow).

To summarize, the reduction of the tropical upwelling, and the local strengthening in the middle latitudes and the decrease in the Arctic of the mid-latitude downwelling compose the latitudinal structural changes of the BDC after 2000. The decline of upwelling is in accord with the BDC weakening. The local acceleration of the mid-latitude downwelling results from the branch narrowing equatorward which is related to weak planetary wave activity and coldthe polar vortex enhancement.

217

218 **3.3 Relationship between BDC and WV**

The decline of the upwelling branch controls the increase of the tropical lower 219 stratospheric WV. According to 'tape recorder' effect, the maximum warming of 1.5K 220 221 (Fig.1e3b) in tropical lowermost stratosphere, as-caused by the decline of -upwelling in the weakening BDC, results in an increase of the tropical lower stratospheric WV 222 by the dehydration of air crossing the cold-point tropical tropopause (Schiller et al., 223 2009) in 2001-2011. GOZCARDS data show that WV at 68hPa in 2001-2011 224 increased by the linear rate of 0.60 ppmv decade⁻¹ ($\frac{17.518}{17.518}$ · decade⁻¹) which 225 exceeds the significance test at the 99% confidence level (Fig.2a4a). WV at 46 hPa 226 showed a similar trend of 0.27 ppmv decade⁻¹ (7.37% · decade⁻¹) that passes 227 statistical significance test at the 90% confidence level (Fig.2b4b). Randel et al. (2006) 228 229 and Dhomse et al. (2008) indicated the opposite variation of WV during several years including 2000, a distinct mutational year (Fueglistaler, 2012). Here we extend the 230 data series and discuss WV during 2001-2011 after the mutation. Fueglistaler (2012) 231 analyzed tropical stratospheric WV from HALOE and also found the trend after 2000. 232 233

234

235 3.4 Relationship between BDC and HCl

236	There is a trend shift of boreal HCl in 2006/2007. GOZCARDS data at 32hPa
237	show that mid-latitude (40°N-60°N) stratospheric HCl had a decreasing trend at a
238	linear rate -0.40 ppbv·decade ⁻¹ (- $\frac{25.526}{\%}$ · decade ⁻¹) before 2006 and had an
239	increasing trend of 0.38 ppbv \cdot decade ⁻¹ ($\frac{24.725}{8}$ % \cdot decade ⁻¹) after 2006
240	(Fig. <u>2e4c</u>). Both trends exceed the significance test at the 99% confidence level. In
241	order to show only the downwelling impacts, HCl regressed against the monthly solar
242	10.7cm flux (at the 99% confidence level) was removed from the monthly HCl data in
243	Fig.2d. Trend shift of the residual HCl remains in 2006/2007. Trend after 2006 is 0.32
244	ppbv-decade ⁻¹ and still passes statistical significance test at the 90% confidence
245	level. HCl-At 46hPa in the mid-latitudes, HCl had similar trends of -0.15 ppbv-
246	decade ⁻¹ (-11.3 <u>11</u> % · decade ⁻¹) before 2006 and 0.33 ppbv· decade ⁻¹ (24.5 <u>25</u> %
247	· decade ⁻¹) after 2006 at 46hPa , -0.44 ppbv· decade⁻¹ (-24.8% · decade⁻¹) before
248	2006 and 0.17 ppbv·decade⁻¹ (9.8% ·decade⁻¹) after 2006 at 22hPa (pictures
249	omitted), which was also noted by Mahieu et al. (2014). However, HCl in the Arctic
250	(70°N-90°N) at 22 hPa continuously decreased after 2001 (Fig.2e4d) - Declining trend
251	is with -0.32 ppbv decade ⁻¹ which passes statistical significance test at the 95%
252	confidence level. At the same time, the minimal HCl concentration in winter reduced
253	faster after 2006 (in the ellipse in Fig.4d). After removal of the solar cycle signals
254	(Fig.2f), HCl reduced faster after 2006/2007 which is opposite to that in the
255	mid-latitudes. It There is a similar declining trend of HCl at 15hPa in the Arctic
256	(pictures omitted).

According to the atmospheric continuity equation, mass flux determines the local 257 HCl concentration. Under the approximate conditions of no divergence and no 258 259 horizontal HCl gradient, HCl transport is controlled by both vertical velocity and the vertical HCl gradient. The downward motion might bring higher concentration HCl 260 from higher altitudes down to the middle stratosphere. HCl vertical gradient in the 261 mid-latitude middle stratosphere is larger than that in high latitudes (in the black 262 rectangle in Fig.5), which may enhance HCl downward transport and hence magnify 263 the HCl positive trend in the middle latitudes caused by the speedup of downwelling 264 in recent years. Similarly, the HCl decrease in the Arctic (in the ellipse in Fig.4d and 265

266 Fig. 4a of Mahieu et al., 2014) after 2006/2007 can support the local slowdown of
267 downwelling.

The positive trend of the age of air for the northern mid-latitudes during the last 268 decade (Ploeger et al., 2015; Mahieu et al., 2014; Stiller et al., 2012) can indicate the 269 weakening of the BDC roughly (Fig.1a) but cannot show the structural change of the 270 BDC (black thin hollow arrows in Fig.3a). The larger vertical gradient located in the 271 mid-latitude middle stratosphere (Fig.5) where downwelling accelerated may be the 272 key point for HCl but not for total air. The weakening of the BDC (not only the 273 tropical upwelling but also the total extratropical downwelling) may increase HCl 274 concentration due to air becoming older and more photolyzed. But HCl decreased 275 near tropical stratopause (source region) during 2001-2011 (Froidevaux et al., 2015) 276 with CFCs decreased, which can exclude the influence of an increasing HCl source by 277 photolysis on the middle stratospheric HCl increase in the mid-latitudes. Therefore, 278 from a dynamic point of view, the positive trend of mid-latitude HCl is better 279 explained by the speedup of local downwelling. the strength of the local downwelling 280 281 dominates the extratropical stratospheric local HCl concentration by downward 282 transport. Therefore, the increase in the mid latitudes and decrease in the Arctic of HCl after 2006/2007 can support the local speedup in mid-latitude and slowdown in 283 the Arctic of the downwelling resulting from the narrowing equatorward of the 284 downwelling branch of BDC. 285

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287

4 Conclusion and Discussion

The trends of <u>the</u> boreal BDC tended to decrease and create latitudinal structural changes after 2000. The latitudinal structure of recent changes in the boreal BDC presented the reduction of the tropical upwelling and the local strengthening of the mid-latitude downwelling. The acceleration of downwelling in the mid-latitudes is caused by the <u>equatorward narrowing of the BDC</u> branch<u>narrowing</u>-equatorward which is related to <u>the a</u> weaker planetary wave activity and <u>the a</u> stronger polar vortex.

295

In addition, the variation of WV in the tropics and HCl in the extratropics support

the latitudinal structural changes of the BDC. The increasing <u>lower</u> tropical-<u>lower</u> stratospheric WV results from the higher tropical CPTT which is mainly caused by the decline of the upwelling after 2000. The increasing HCl in the mid-latitudes is caused by the local speedup of downwelling after 2006/2007.

Numerous simulation studies (Eichelberger et al., 2005; Garcia and Randel, 2008; 300 301 Butchart et al., 2010; Li et al., 2010; Fleming et al., 2011; Douglass et al., 2014) claimed that BDC will be stronger in the future. Those results are not consistent with 302 the observations described in this paper. However, Winter and Bourqui (2011) found 303 that mid-latitude zonal symmetric surface heating forcing, decreasing the meridional 304 temperature gradient between the tropics and the mid-latitudes, results in a decline of 305 the upward propagation of PWs in the polar region and enhances the polar vortex. 306 Realistically, zonally asymmetric land ocean warming is more likely to affect the 307 308 PWs activity. Hu et al. (2014) confirmed that increasing the meridional temperature gradient of sea surface temperature (SST) could enhance the BDC and would be 309 accompanied by increased PWs activity and a decline in the polar vortex. But in 310 311 2001-2011, the Pacific Decadal Oscillation (PDO) shifted from the warm phase to the cold phase. An increase of North Pacific SST weakened the meridional temperature 312 gradient between the tropics and the mid-latitudes which was opposite to the 313 simulation conditions of Hu et al. (2014). It appears the modification of the PDO in 314 the meridional temperature gradient is a likely reason for the weaker PWs activity and 315 the latitudinal structural changes in the boreal BDC in the beginning of the 21st 316 century. 317

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319 Acknowledgements

320 The provisions of online data by NASA and ECMWF are gratefully acknowledged. .

321 This study was jointly supported by National Natural Science Foundation of China

322 (Grant No. 41375047, 41305039, 91537213), China Scholarship Fund and Priority

323 Academic Program Development of Jiangsu Higher Education Institutions (PAPD).

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416 Figures



Fig.1 (a) Time series of 45°N-75°N mean vertical EPF from ERA-Interim in DJF at
50hPa, red lines show piecewise linear fits; (b) 10°S-10°N mean monthly temperature
from GOZCARDS in DJFM at 68hPa, red thin line shows piecewise linear fit and the
gray thick curve shows the 25-month running mean;





Fig.2 (a) Mean EPF (black thin arrows) of PWs and zonal wind (blue contours, m s⁻¹)
in DJF of 2001-2011 from ERA-Interim, red thick hollow arrow indicates the mean
BDC; (b) as in (a), but for the anomalies between 2007-2011 and 2001-2005;





Fig.3 (a) Mean TEM vertical velocity $(10^{-4} \text{ Pa s}^{-1}, \text{ shading})$ in DJF of 2001-2011, the 436 anomalies of vertical velocity (10⁻⁴ Pa s⁻¹, black contours) and zonal wind (blue 437 contours, m s⁻¹) in DJF between 2007-2011 and 2001-2005, red thick hollow arrow 438 indicates the mean BDC, black thin hollow arrows indicate the local BDC anomalies; 439 (b) single variable linear regression of temperature (in DJFM, unit, Kelvin) on 440 45°N-75°N mean EPFp (in DJF at 50 hPa) in 2001-2011, Stippled regions passed the 441 90% (filled) and 85% (hollow) confidence test levels; (c) 35°N-65°N TEM vertical 442 speed of BDC in DJF at 20hPa, black and red curves respectively indicate 3 and 5 443 year running means; (d) as in (c), but at 65°N-85°N. 444

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Fig.4 Time series of monthly H2O and HCl from GOZCARDS, thick gray curve for
25-month running mean and thin red line for piecewise linear fit; (a, b) H2O in tropics
10°S-10°N at 68hPa and 46hPa; (c) HCl in mid-latitude 40°N-60°N at 32hPa; (d) HCl
in the Arctic 70°N-90°N at 22hPa;



458 Fig.5 HCl profiles in DJF with different latitudes from MLS, decreasing vertical459 gradient with latitude in black rectangle.