

1 **New fire diurnal cycle characterizations to improve fire radiative**
2 **energy assessments made from MODIS observations**

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15

16 **Abstract**

17

18 Accurate near real time fire emissions estimates are required for air quality forecasts. To date, most
19 approaches are based on satellite-derived estimates of fire radiative power (FRP), which can be
20 converted to fire radiative energy (FRE) which is directly related to fire emissions. Uncertainties in
21 these FRE estimations are often substantial. This is for a large part because the most often used low-
22 Earth orbit satellite-based instruments such as the MODerate-resolution Imaging Spectroradiometer
23 (MODIS) have a relatively poor sampling of the usually pronounced fire diurnal cycle. In this paper we
24 explore the spatial variation of this fire diurnal cycle and its drivers using data from the geostationary
25 Meteosat Spinning Enhanced Visible and Infrared Imager (SEVIRI). In addition, we sampled data from
26 the SEVIRI instrument at MODIS detection opportunities to develop two approaches to estimate
27 hourly FRE based on MODIS active fire detections. The first approach ignored the fire diurnal cycle,
28 assuming persistent fire activity between two MODIS observations, while the second approach
29 combined knowledge on the climatology of the fire diurnal cycle with active fire detections to
30 estimate hourly FRE. The full SEVIRI time-series, providing full coverage of the fire diurnal cycle, were
31 used to evaluate the results. Our study period comprised of three years (2010–2012), and we
32 focussed on Africa and the Mediterranean basin to avoid the use of potentially lower quality SEVIRI
33 data obtained at very far off-nadir view angles. We found that the fire diurnal cycle varies
34 substantially over the study region, and depends on both fuel and weather conditions. For example,
35 more “intense” fires characterized by a fire diurnal cycle with high peak fire activity, long duration
36 over the day, and with nighttime fire activity are most common in areas of large fire size (i.e., large
37 burned area per fire event). These areas are most prevalent in relatively arid regions. Ignoring the
38 fire diurnal cycle generally resulted in an overestimation of FRE, while including information on the
39 climatology of the fire diurnal cycle improved FRE estimates. The approach based on knowledge of
40 the climatology of the fire diurnal cycle also improved distribution of FRE over the day, although only
41 when aggregating model results to coarser spatial and/or temporal scale good correlation was found
42 with the full SEVIRI hourly reference dataset. We recommend the use of regionally varying fire
43 diurnal cycle information within the Global Fire Assimilation System (GFAS) used in the Copernicus
44 Atmosphere Monitoring Services, which will improve FRE estimates and may allow for further
45 reconciliation of biomass burning emission estimates from different inventories.

46

47 1 Introduction

48

49 Landscape fires are a global phenomena, and the annually burned area is roughly equivalent to the
50 area of India (Giglio et al., 2013). Most burned area occurs in the savannas of Africa, Australia, and
51 South America, where they shape ecosystem dynamics and due to their scale are an important
52 source of global emissions of (greenhouse) gases and aerosols (Seiler and Crutzen, 1980; Bowman et
53 al., 2009). Fires affect air quality both locally and regionally (Langmann et al., 2009), with recent
54 studies putting mortality rates over 300000 annually due to exposure to smoke (Johnston et al.,
55 2012).

56

57 Traditionally, the amount of dry matter burned and quantity of trace gases and aerosols emitted
58 have been calculated using biome-specific fire return intervals and estimates of the total amount of
59 biomass as well as the fraction of biomass burned, the combustion completeness (Seiler and Crutzen,
60 1980). Thanks to new satellite input streams that better capture the spatial and temporal diffuse
61 nature of fires, the estimated fire return intervals have been replaced by direct estimates of monthly,
62 weekly or even daily area burned (Roy et al., 2005; Giglio et al., 2009). In addition, satellite
63 information and biogeochemical modelling have been used to improve estimates of biomass and
64 combustion completeness. However, uncertainties in these bottom-up fire emission estimates are
65 still substantial (Reid et al., 2009; Zhang et al., 2012; Larkin et al., 2014), and they are generally
66 inappropriate for use in near real-time systems partly because the burned area signature is only
67 detectable days to weeks after the actual fire occurrence.

68

69 Hot spot observations from satellite have been used as a proxy for burned area and emissions fluxes
70 in near real time (Freitas et al., 2005; Reid et al., 2009; Wiedinmyer et al., 2011). Another promising
71 and relatively new bottom up approach uses estimates of fire radiative power (FRP) observed from
72 satellite to calculate daily fire radiative energy (FRE). Wooster et al. (2005) found that these FRE
73 estimates scale directly to dry matter burned, potentially circumventing the uncertainties associated
74 with estimating area burned, fuel loads, and the combustion completeness. In addition, FRP
75 observations can be observed and processed in near real time (Xu et al., 2010; Kaiser et al., 2012;
76 Zhang et al., 2012) and can be measured for small fires that remain undetected in burned area
77 products (Roberts et al., 2011; Randerson et al., 2012).

78

79 Hot spot and FRP observations are currently the only available options when modelling exercises
80 require near real time observations, for example in chemical weather forecasts used to predict air
81 quality. The Global Fire Assimilation System (GFAS; Kaiser et al., 2012), for example, is used to
82 estimate global near real time daily fire emissions within the EU-funded project Monitoring
83 Atmospheric Composition and Climate III (MACC-III). GFAS is currently using fire observations from
84 the polar orbiting MODERate-resolution Imaging Spectroradiometer (MODIS) instruments aboard the
85 Terra and Aqua satellites (Giglio et al., 2006). Due to their relative proximity to the Earth, the Terra
86 and Aqua MODIS instruments have a high sensitivity to even quite low FRP (smaller and/or lower
87 intensity) fires. However, they only provide four daily observations under ideal conditions but less
88 when optically thick clouds are present, which may not be enough to fully characterize how fire
89 activity varies over the course of the day. Observations with a much higher temporal resolution are
90 available from geostationary satellites. However, as a consequence of their geostationary position,

91 these satellites individually do not provide global data and are located at greater distance from the
92 Earth resulting in typically coarser pixel sizes than polar orbiting instruments. Since the threshold of
93 detectability of a fire is not only dependent on the instrument but also a function of the pixel area,
94 geostationary sensors have a higher minimum FRP detection limit (typically > 40 MW) than MODIS (~
95 8 MW). They therefore do not observe the lowest FRP component of the fire regime (Roberts et al.,
96 2005; Freeborn et al., 2014).

97
98 Previous studies found that fire activity shows a strong diurnal cycle, and one that is both temporally
99 and spatially variable (Prins and Menzel, 1992; Giglio, 2007; Roberts et al., 2009). The ideal set-up to
100 detect fires would be a high temporal resolution imaging system, sensitive to even the lowest FRP
101 fires, and providing global coverage, but due to the limitations of the orbital characteristics outlined
102 above there is no single platform available to provide this. Therefore the estimation of FRE at a global
103 scale is difficult, with polar orbiting satellites lacking observations to accurately represent the fire
104 diurnal cycle and geostationary satellites being limited to certain regions of the globe and omitting
105 the (rather common) low FRP fires. However, previous studies have developed approaches to
106 estimate FRE based on the combination of data from different satellite systems (Boschetti and Roy,
107 2009; Ellicott et al., 2009; Freeborn et al., 2009, 2011; Vermote et al., 2009).

108
109 Some of these mixed approaches used both burned area and active fire data (Boschetti and Roy,
110 2009; Roberts et al., 2011), which may provide benefits in terms of more accurate FRE determination
111 but cannot be used easily in near real time systems because of the latency in burned area
112 observations. Alternatively, FRP observations of polar orbiting and geostationary satellites can be
113 blended to combine the sensitivity of the MODIS instruments to lower FRP fires and the diurnal
114 sampling characteristics of SEVIRI. Freeborn et al. (2009) developed a database for matching SEVIRI
115 and MODIS FRP observations based on frequency-magnitude statistics, but the samples had to be
116 accumulated over significant spatial areas (5°) to provide matchable statistics, which is incompatible
117 with the need to develop a method operating at high spatial resolution. Freeborn et al. (2011) later
118 presented an alternative approach, estimating FRE using MODIS data accumulated over 8 day
119 periods over which MODIS samples a location at the fullest range of view zenith angles. The
120 relationship between the “true” FRE and the limited number of FRP samples provided by MODIS was
121 derived using SEVIRI FRP time-series sampled at the MODIS sampling interval. Vermote et al. (2009)
122 and Ellicott et al. (2009) used a different approach to create FRE data from MODIS, showing that for
123 several regions of the globe the fire diurnal cycle can be described by a Gaussian function, and used
124 monthly MODIS data to fit the parameters of the Gaussian. Using this approach, a first global
125 estimation of monthly FRE was made (Ellicott et al., 2009). Despite the success of these latter
126 approaches with regard to estimating FRE from MODIS, they are not a solution to the problem posed
127 herein because they require 8 days of consecutive MODIS data and therefore cannot be applied in a
128 near real-time approach.

129
130 Global fire emissions estimates at high spatial and temporal resolutions, ideally produced in near real
131 time, are required to feed into atmospheric models which are under continuous development and
132 run at improved resolutions thanks to increased computational power (Zhang et al., 2012). Higher
133 temporal resolution may also help to reconcile bottom up and top down emission estimates (Mu et
134 al., 2011). None of the approaches mentioned above are, however, suitable for providing this. Due to
135 these limitations current state of the art global near real time emission inventories still ignore

136 possible effects of fire diurnal cycle on their emission estimates (e.g., Wiedinmyer et al., 2011; Kaiser
137 et al., 2012) and may therefore be structurally biased due to the fire diurnal cycle and the MODIS
138 sampling design (e.g., Ichoku et al., 2008; Ellicott et al., 2009; Freeborn et al., 2011).

139

140 The purpose of the work presented here is to better understand the fire diurnal cycle and its
141 spatiotemporal dynamics, in order to develop a new way to include this into a near real time fire
142 emissions estimation framework. First, the spatial distribution and dependencies of the fire diurnal
143 cycle and their effect on active fire detections at MODIS overpasses were explored. Then, data
144 assimilation was used to compare two different methods to derive hourly FRE estimates at 0.1°
145 resolution based on low Earth-orbiting MODIS observations. The first method ignored the fire diurnal
146 cycle, and was used as a reference to better understand the combined effect of the fire diurnal cycle
147 and the MODIS sampling design on hourly FRE estimates. The second method combined knowledge
148 on the fire diurnal cycle with active fire detections at MODIS overpasses. Following previous studies
149 (Freeborn et al., 2009, 2011), we used FRP observations derived from data collected by the
150 geostationary SEVIRI instrument at MODIS detection opportunities, rather than actual MODIS
151 observations, to drive the two model approaches and we evaluated the model results against the full
152 SEVIRI time-series. We used three years of active fire data (2010–2012) across Africa and the
153 Mediterranean basin to include a wide range of climates and land cover types, and avoid the use of
154 SEVIRI observations obtained at very far off-nadir angles over South America and northern Europe
155 (Freeborn et al., 2014). Results are intended for application in GFAS within EU’s Copernicus
156 Atmosphere Monitoring Service (CAMS, <http://atmosphere.copernicus.eu>).

157

158 **2 Data**

159

160 To explore the spatiotemporal dynamics of the fire diurnal cycle, we used hourly temporal resolution
161 FRP data derived from 15 min observations made by the SEVIRI instrument hosted onboard the
162 geostationary Meteosat satellite (Sect. 2.1). However, to drive the models developed here we only
163 used SEVIRI FRP observations made at the overpass times of the MODIS polar orbiting sensors (Sect.
164 2.2), whilst the hourly temporal resolution SEVIRI time-series were used to evaluate the results. Land
165 cover characteristics (Sect. 2.3), along with data on fire size (Sect. 2.4), were used to better
166 understand the spatial distribution of fire diurnal cycle. These datasets are described in more detail
167 below, followed by the methods used in Sect. 3.

168

169 ***2.1 SEVIRI fire radiative power (FRP)***

170

171 The SEVIRI instrument aboard the geostationary Meteosat Second Generation (MSG) series of
172 satellites provides coverage of the full Earth disk every 15 min in 12 spectral bands (Schmetz et al.,
173 2002). The Meteosat SEVIRI FRP-PIXEL product contains per-pixel fire radiative power data along with
174 FRP uncertainty metrics produced from the full spatial and temporal resolution SEVIRI observations
175 (Wooster et al., 2015). The FRP-PIXEL product is produced using an operational version of the
176 geostationary Fire Thermal Anomaly (FTA) algorithm described in Roberts and Wooster (2008), and
177 the product and its performance characteristics are described in Wooster et al. (2015). The FRP-PIXEL

178 products are freely available from the Land Surface Analysis Satellite Applications Facility (LSASAF;
179 <http://landsaf.meteo.pt>), from the EUMETSAT EO Portal (<https://eoportal.eumetsat.int>) or via the
180 EUMETCAST dissemination service (<http://www.eumetsat.int>) in both real-time and archived form,
181 as detailed in Wooster et al. (2015). The Meteosat satellites are located at 0° longitude and latitude,
182 and at nadir the SEVIRI pixels cover 3 km x 3 km on the ground, but this degrades with increasing
183 view angle away from the West African sub-satellite point (Freeborn et al., 2011; Roberts et al.,
184 2015). The FRP-PIXEL product data used here were obtained from the LSA SAF and were rescaled to
185 an hourly 0.1° resolution with the GFAS gridding algorithm explained in Kaiser et al. (2012). Missing
186 FRP values in individual observations within the hour (e.g., due to smoke or short periods of cloud
187 cover) were thus implicitly ignored. A single 0.1° grid cell comprises over 13 SEVIRI pixels close to the
188 sub-satellite point (equatorial West Africa) and this reduces to around 6 SEVIRI pixels at greater of
189 nadir angles (e.g., Portugal and Madagascar). Data were archived in the Meteorological Archival and
190 Retrieval System (MARS) of the European Centre for Medium range Weather Forecasting (ECMWF)
191 prior to their use herein.

192

193 **2.2 MODIS detection opportunity**

194

195 The two MODIS sensors on board of the Terra and Aqua satellites provide 4 daily overpasses in most
196 Earth locations, albeit sometimes at view angles in excess of 45° where the product performance is
197 somewhat degraded (Freeborn et al., 2011). At nadir the MODIS thermal channel spatial resolution is
198 1 km, but decreases away from the swath centre (Freeborn et al., 2011). We used the MODIS MOD03
199 (Terra) and MYD03 (Aqua) geolocation products to determine where and when MODIS data were
200 collected within the SEVIRI Earth disk. As cloud cover may further limit the fire detection opportunity,
201 we used the data quality and cloud cover information of the MOD14 and MYD14 active fire products
202 to filter out grid cells with cloud cover (Giglio et al., 2006). Here we define the detection opportunity
203 as the ability to make unobstructed observations, and the MODIS detection opportunity was derived
204 by combining the MOD03, MYD03, MOD14 and MYD14 products, combining overpass times and
205 cloud cover. We used MODIS data from Collection 5. Like the SEVIRI data, these data were rescaled
206 to hourly 0.1° resolution with the GFAS gridding algorithm and archived in MARS (Kaiser et al., 2012).
207 The data were archived for the Terra and Aqua satellites separately. The original MODIS swath data
208 can be downloaded from NASA at <http://reverb.echo.nasa.gov>.

209

210 **2.3 MODIS Land cover**

211

212 The dominant land cover type was derived from the MODIS MCD12C1 land cover product, which
213 provides 0.05° spatial resolution annual information on land cover (Friedl et al., 2002). We calculated
214 the dominant land cover type for each grid cell as the land cover type that on average covered the
215 largest fraction during the study period (2010–2012). The University of Maryland (UMD) classification
216 scheme was used, and data was rescaled to 0.1° resolution. Because we only considered Africa and
217 the Mediterranean basin in this study, and because in some land cover classes very few fires
218 occurred, we could merge some land cover classes that were of relatively little importance for our
219 study. Specifically, all forest classes within the tropics were binned into the tropical forest class, while

220 extratropical forests were all labelled temperate forest. Open and closed shrublands were merged
221 into one shrubland class, and urban and built-up, barren or sparsely vegetated into grasslands.
222

223 **2.4 Fire size**

224
225 Here we define the fire size for a certain grid cell as the mean burned area per fire event, weighted
226 by their total area burnt (when calculating the mean, a fire event burning 100 km² is assigned one
227 hundred times the weight of an event burning 1 km²). The MODIS MCD64A1 burned area product
228 provides daily mapped estimates of global burned area (Giglio et al., 2009). We applied the methods
229 described by Archibald and Roy (2009) to derive a global mean fire “size” (area) map using data over
230 our study period (2010–2012). We made one modification to the approach described by Archibald
231 and Roy (2009): we assumed that two neighbouring burned area grid cells only belonged to the same
232 fire if the burn date was no longer than two days apart (instead of 8 days). We believe that overall
233 this provides a better estimation of the fire size in our study region, as the vast majority of fires here
234 are grass fires, occurring outside tropical forest zones and thus spreading relatively fast while being
235 relatively less often obstructed by cloud cover. Consequently, the uncertainty in burn date is
236 generally low in our study region (Giglio et al., 2013) and so the two day thresholds was deemed
237 more appropriate.
238

239 **3 Methods**

240
241 Our overall goal within GFAS is to provide hourly estimates of FRE at 0.1° spatial resolution, based on
242 the limited number of MODIS overpasses available each day at each grid cell location. This limited
243 number of daily MODIS observations, in combination with the often pronounced fire diurnal cycle,
244 are the major obstacles in providing the required output. We first studied the spatiotemporal
245 variation of the fire diurnal cycle, in an attempt to understand its variability (Sect. 3.1). Then, we
246 explored the way the fire diurnal cycle affects active fire detections made at the MODIS sampling
247 times (Sect. 3.2). Using this knowledge we explored a new method to parameterize the fire diurnal
248 cycle, and compared results to a modelling approach in which the fire diurnal cycle is ignored.
249 Building on the work of Freeborn et al. (2009, 2011), to drive the modelling approaches we used
250 SEVIRI data sampled at the MODIS detection opportunities (according to the hourly data
251 representation introduced above), rather than actual MODIS observations (Sect. 3.2). This allowed us
252 to focus on the issue of diurnal cycle sampling rather than simultaneously dealing with the issue of
253 MODIS and SEVIRI’s differential sensitivity to active fires (Freeborn et al., 2009).
254

255 Using data assimilation we combined the discrete actual SEVIRI observations, made at the time of the
256 MODIS detection opportunities, with hourly predictions of fire activity – using their combination to
257 create continuous hourly best estimate FRE time-series (Sect. 3.3). We developed two prediction
258 methods. The first method assumed persistent fire activity until the next satellite detection
259 opportunity, and provides further insights into the combined effect of the fire diurnal cycle and the
260 MODIS sampling design on hourly FRE estimates when the fire diurnal cycle is ignored (Sect. 3.4). The
261 second method followed previous studies and used a Gaussian function to predict fire development

262 over the day (Vermote et al., 2009). By combining prior knowledge about the climatology of the fire
 263 diurnal cycle with active fire observations at MODIS overpasses to estimate the parameters of the
 264 Gaussian function, this approach provides a possible pathway to implement the fire diurnal cycle into
 265 the near real time fire emission modelling framework (Sect. 3.5). Comparing the results of the two
 266 approaches to those from the full hourly SEVIRI time-series allowed us to determine and discuss their
 267 strengths and limitations (Sect. 3.6).
 268

269 **3.1 Exploring the fire diurnal cycle**

270
 271 We started exploring the fire diurnal cycle and its drivers. A Gaussian function was optimally fitted
 272 (least squares) to the hourly SEVIRI observations $\tilde{\rho}_{SEV(t)}$ for each grid cell and day of fire activity
 273 during the study period:
 274

$$\tilde{\rho}_{SEV(t)} = \rho_{base} + (\rho_{peak} - \rho_{base})e^{-\frac{(h_t - h_{peak})^2}{2\sigma^2}}. \quad (1)$$

275
 276 Where ρ_{base} corresponds to the nighttime fire activity, ρ_{peak} to the maximum FRP for a given day, σ
 277 is the standard deviation (SD) of FRE distribution over the day (dependent on fire duration), h_t is the
 278 local solar time at time step t and h_{peak} is the local hour at which FRP reaches its daily maximum.
 279 This resulted in a database containing hourly time-series of $\tilde{\rho}_{SEV(t)}$ and the fitted Gaussian function,
 280 and daily time-series of optimal parameter values of the Gaussian function for each grid cell. At the
 281 same time we also kept track of hourly MODIS detection opportunities. This enabled us get a better
 282 understanding of structural errors caused by the MODIS sampling design in relation to the actual fire
 283 diurnal cycle.
 284

285 Although the fire diurnal cycle as observed by SEVIRI is comparable to that which would be observed
 286 by MODIS if it had the same temporal sampling ability, it is a little different due to SEVIRI's inability to
 287 discriminate the lowest FRP fire pixels which typically dominate more towards the start and end of
 288 the daily fire cycle, but which are also present along with often higher FRP pixels towards the diurnal
 289 cycle maxima (Freeborn et al., 2009). To gauge the magnitude of the effect Freeborn et al. (2009)
 290 derived the "virtual MODIS" fire product that has the temporal sampling of SEVIRI and the sensitivity
 291 to fire of MODIS. They found that the full-width at half maximum height (i.e., the width of the diurnal
 292 cycle at half of the daily FRP maximum value) of the diurnal cycles derived from the SEVIRI and the
 293 "virtual MODIS" datasets are very similar, it is the amplitude and the full-width at base height of the
 294 two cycles, which are more different. In terms of total FRE emitted, the latter is of less importance,
 295 here we followed Freeborn et al. (2011) in assuming that the diurnal cycles from SEVIRI and MODIS
 296 are sufficiently similar.
 297

298 In order to visualize the spatial distribution of the fire diurnal cycle, the climatological diurnal cycle
 299 was calculated for each grid cell depending on the mean parameter values of the Gaussian function
 300 weighted for daily FRE, including all days of fire activity during the study period without cloud
 301 obscurance. To get a better understanding of the drivers of the fire diurnal cycle these results were

302 compared to land cover and aspects of the fire regime (fire size, total annual FRE, and the annual
303 number of days with fire activity), see Sect. 2.
304

305 **3.2 Sampling SEVIRI data at MODIS detection opportunities**

306
307 During the data assimilation, SEVIRI observations at MODIS detection opportunities were used to
308 drive the models. Here, SEVIRI observations for a given hour t are given by $\tilde{\rho}_{SEV(t)}$ and SEVIRI fraction
309 of observed area by $\tilde{\alpha}_{SEV(t)}$; in the same way, observations of the MODIS instruments are given by
310 $\tilde{\rho}_{MOD(t)}$ and $\tilde{\alpha}_{MOD(t)}$. Therefore input for the models, i.e., the SEVIRI observations at MODIS
311 detection opportunity times ($\tilde{\rho}_t$ and $\tilde{\alpha}_t$) for a given hour t are given by:
312

$$\tilde{\rho}_t = \tilde{\rho}_{SEV(t)} \quad (2)$$

$$\tilde{\alpha}_t = \tilde{\alpha}_{MOD(t)}. \quad (3)$$

313
314 For clarity, we assumed that the observed FRP $\tilde{\rho}_t$ is zero when there was no MODIS detection
315 opportunity. Anyhow, during the data assimilation $\tilde{\rho}_t$ was weighted for observed area $\tilde{\alpha}_t$, which was
316 zero when there was no observation.

317
318 SEVIRI data sampled at MODIS detection opportunities were compared to the full SEVIRI hourly time-
319 series to explore the effect of the fire diurnal cycle on the daily sampling at MODIS overpasses. More
320 specifically we calculated the percentage of FRE emitted on days without any active fire detection at
321 MODIS detection opportunities, and the total daily number of MODIS overpasses during the fire
322 season. The latter was calculated by weighing the mean number of monthly detection opportunities
323 at MODIS overpasses by monthly total detected FRP, thus giving the largest weight to the month with
324 most fire activity (ignoring cloud cover).
325

326 **3.3 Data assimilation**

327
328 We used a modified version the fire data assimilation methodology of GFAS to allow representation
329 of the fire diurnal cycle. GFAS assumes that the availability of observations dominates the error
330 budget of the global FRP fields. It approximates these errors by further assuming the FRP variance to
331 be inversely proportional to the fraction of observed area $\tilde{\alpha}_t$. Thus the variance increases with
332 decreasing partial cloud cover and with the number of satellite observations. In the following data
333 assimilation, GFAS fills observation gaps with a Kalman filter, in which current observations are
334 combined with information from earlier ones. The Kalman filter has a time step of 1 day. It uses a
335 trivial predictive model for the temporal evolution of FRP (i.e., persistence), and assumes for the
336 accuracy of the 1 day FRP prediction that the variance increases by a factor of 9 (Kaiser et al., 2012).
337

338 Our modifications affected the step size and the FRP prediction model. The former was set to 1h to
339 be able to represent a diurnal cycle. For calculating the FRP prediction $\check{\rho}_t$, we investigated two
340 different approaches (Sects. 3.4 and 3.5). In both cases, we assumed for the accuracy of the 1h FRP
341 prediction that the variance increases by a factor of 4. Lowering the value compared to the daily

342 GFAS is motivated by the shorter time step used in our study. However, lowering it too much would
 343 not give sufficient weight to new FRP observations. Thus the analysis FRP $\hat{\rho}_t$ and “fraction of
 344 observed area” $\hat{\alpha}_t$ were calculated at each 1h time step by optimal interpolation as follows, cf. Eqs.
 345 (32)–(33) of Kaiser et al. (2012):

$$\hat{\rho}_t = \frac{1}{\hat{\alpha}_t} \left(\frac{\hat{\alpha}_{t-1}}{5} \check{\rho}_t + \tilde{\alpha}_t \tilde{\rho}_t \right) \quad (4)$$

347
 348 with $\check{\rho}_t$ according to Sects. 3.4 and 3.5 and

$$\hat{\alpha}_t = \frac{\hat{\alpha}_{t-1}}{5} + \tilde{\alpha}_t. \quad (5)$$

349

350 **3.4 Persistent approach**

351

352 Applying the daily persistence approach of Kaiser et al. (2012) to hourly time resolution, we first
 353 explored the most parsimonious approach that predicts FRP $\check{\rho}_t$ as being equal to the FRP of the
 354 previous time step’s analysis:

355

$$\check{\rho}_t = \hat{\rho}_{t-1}. \quad (6)$$

356

357 This approach provided insights in the spatiotemporal consequences for FRE estimation when
 358 information on the fire diurnal cycle is not incorporated.

359

360 **3.5 Climatological approach**

361

362 In the second approach we followed previous studies of Vermote et al. (2009) and Ellicot et al. (2009)
 363 and the recommendation in Kaiser et al. (2009) to use a Gaussian function to describe a “standard
 364 fire diurnal cycle”. Wooster et al. (2005) and Roberts et al. (2009) already demonstrated that SEVIRI
 365 observations sample the diurnal cycle of large fires well, and for some individual large fires show FRP
 366 time-series that depict diurnal characteristics appearing close to Gaussian in nature even at 15 min
 367 temporal resolution. The prediction was calculated by optimally fitting a Gaussian function through
 368 the last 24h of analysis:

369

$$\check{\rho}_t = \rho_{base} + (\rho_{peak} - \rho_{base}) e^{-\frac{(h_t - h_{peak})^2}{2\sigma^2}} \quad (7)$$

370

371 However, only h_{peak} was optimally fitted, by minimizing the sum of least squares between the
 372 Gaussian function and the previous 24h of the analysis:

373

$$\hat{\rho}_{t-24}, \hat{\rho}_{t-23}, \dots, \hat{\rho}_{t-1}. \quad (8)$$

374

375 Following previous studies that found that fire diurnal cycle is land cover dependent (Giglio, 2007;
376 Roberts et al., 2009; Vermote et al., 2009; Freeborn et al., 2011), we used land cover (LC) average
377 values σ_{LC} for σ (weighted by FRE). Values of ρ_{base} and ρ_{peak} on the other hand could be directly
378 related to daily MODIS observations. We followed Vermote et al. (2009) to use the mean of the
379 nighttime (defined here as 6p.m.–6a.m. the next day) observations at MODIS detection opportunities
380 to determine ρ_{base} . To relate SEVIRI observations at MODIS detection opportunities to ρ_{peak} the ratio
381 of mean daytime (6a.m.–6p.m.) FRP observations at MODIS detection opportunities to mean ρ_{peak}
382 was calculated per land cover type. We used per land cover average values for scaling the daytime
383 observations at MODIS detection opportunities to ρ_{peak} rather than the values found per grid cell to
384 keep the model generic and globally applicable. Finally, if there were no active fires observed during
385 the previous 24h, we forced the prediction to be zero, to prevent fires from continuing during long
386 periods of no observations.
387

388 **3.6 Model evaluation**

389
390 The estimated hourly FRE fields (or analysis; $\hat{\rho}_t$) resulting from the two modelling approaches
391 (persistent and climatological) were evaluated via comparison to those derived from the hourly
392 SEVIRI time-series (see Sect. 2.1). Two criteria were used to evaluate model performance: first, the
393 spatial distribution of FRE estimates; and second, the temporal distribution of FRE. The spatial
394 performance of the modelling approaches was assessed via their ability to reproduce the annual
395 mean FRE per land cover type, and by comparing the spatial distribution of FRE as estimated by the
396 modelling approaches and as derived from SEVIRI over the study region and period. The temporal
397 performance was assessed via the ability of the model to allocate the emitted energy in the right grid
398 cell at the right moment in time. Here we used Pearson's r between the modelled and observed
399 (SEVIRI) FRE time-series at four different spatiotemporal resolutions (0.1° and 1° spatial, and hourly
400 and daily temporal resolution). Each spatiotemporal scale provides unique information on the model
401 performance. Correlation coefficients at hourly resolution depend on the ability of the model to
402 estimate the distribution of fire activity over the day, while daily aggregated estimates provide
403 insights in the ability to get overall budgets right. In a similar way the two spatial resolutions provide
404 information on the ability of the model to resolve high resolution distribution of fire activity and
405 more regional model performance. When calculating Pearson's r between the hourly model results
406 and SEVIRI data we included cloud free days only, while the daily model results were compared to
407 the full cloud cover corrected SEVIRI times series, using a simple cloud cover correction method
408 explained below.

409
410 Finally, we compared daily regional aggregated FRE time-series for several study regions of the two
411 modelling approaches and SEVIRI. In order to compare daily regional time-series to the model, a
412 cloud cover correction needed to be carried out. Since persistent cloud cover is relatively rare during
413 the burning season in most parts of Africa, we chose a simple gap filling approach where the value of
414 the last cloud-free observation is assumed to be valid until the next cloud-free observation, which is
415 consistent with the observation gap filling in the daily GFAS.

416

417 4 Results

418

419 4.1 The diurnal cycle and MODIS sampling

420

421 First, we present the results related to the spatial distribution of the fire diurnal cycle, and assess the
422 impact of the fire diurnal cycle on active fire observations made at the times of the MODIS overpass.
423 The spatial distribution of the fire diurnal cycle was explored by optimally fitting a Gaussian function
424 to the hourly, 0.1° SEVIRI FRP time-series. Reasonable overall correlations between SEVIRI and the
425 optimally fitted Gaussian functions were found (Pearson's $r = 0.55$; weighted mean for all grid cells),
426 while a Gaussian was better able to describe hourly fire activity in regions where fires could spread
427 over large areas and were characterized by high ρ_{peak} (e.g., for fire size $< 10 \text{ km}^2$ $r = 0.51$, for $10\text{--}50$
428 km^2 $r = 0.56$, and $> 50 \text{ km}^2$ $r = 0.63$). This is likely to be in part related to the fact that characterisation
429 of the diurnal cycle of "smaller" fires will be more affected by instances of SEVIRI failing to detect one
430 or more of their fire pixels than would larger fires, hence introducing more variability into the
431 apparent diurnal cycle. Whilst the SEVIRI FRP-PIXEL product shows apparently the best performance
432 metrics of any current geostationary fire product derived from SEVIRI data (Baldassarre et al., 2015),
433 such failures in active fire pixel detection clearly occur, for example simply due to fire pixels being
434 too low in their FRP to detect by SEVIRI, along with a variety of potential other factors (Wooster et
435 al., 2015).

436

437 Figure 1 shows an example of two 0.1° grid cells in which the hourly average FRP maxima reached
438 relatively high levels, well in excess of 1 GW, and fire persisted for several days. As with the individual
439 fires, shown by SEVIRI in Wooster et al. (2005) and Roberts et al. (2009), the FRP from these fires
440 appears to drop to zero, or near zero, every night. This is a consequence both of the actual FRP from
441 the fire significantly diminishing at this time due to, for example, fuel moisture, wind and other
442 ambient atmospheric conditions being far less conducive to intense fire activity by night than by day
443 (Hély et al., 2003; Gambiza et al., 2005), but also because some fire pixels will have FRPs below the
444 SEVIRI active fire pixel detection limit of around 40 MW (Roberts and Wooster, 2008). At the start of
445 the following day, fuel moisture and ambient atmospheric conditions generally become more
446 conducive to fire, and fire intensities and rates of spread typically increase once more such that more
447 of the fire-affected pixels breach the SEVIRI FRP detection limit (Roberts et al., 2009).

448

449 The results shown in Fig. 1 indicate that high FRP, relatively long-lived fire activity is rather well
450 described by a Gaussian function, even at this 0.1°, hourly resolution which is significantly higher
451 than that used in previous studies fitting Gaussian descriptors to remotely sensed measures of active
452 fire activity. At the same time, it also became apparent that observations from a MODIS-type
453 sampling interval are not always representative of the daily fire activity. The inability of the MODIS
454 sampling times to provide representative observations is well illustrated in Fig. 1a, where on the first
455 day of the fire the morning and afternoon time of MODIS sampling slot almost completely missed the
456 fire activity.

457

458 The shape of the Gaussian function, and consequently the parameters: SD (σ) peak fire activity (ρ_{peak})
459 and corresponding hour (h_{peak}), varied considerably over the individual days (Fig. 1). For example, in

460 the African savanna grid cell (Fig. 1c), fire activity on day 3 continued longer in the afternoon
461 compared to day 4, when conditions some-how became less favourable for maintaining the fire
462 earlier in the afternoon. Therefore, the shape of the fire diurnal cycle is dependent on
463 spatiotemporal scale. When diurnal fire activity was aggregated over several days, which can be
464 compared to using a coarser temporal or spatial resolution, increased as compared to fire activity for
465 individual days (compare Fig. 1a with b, and Fig. 1c with d). The relatively narrow diurnal cycle of the
466 individual days have varying peak hours of fire activity, so that the sum of it is wider than any of the
467 individual cycles and the peak fire activity less pronounced.

468
469 In addition to an observed variability in the fire diurnal cycle seen on different days, we found
470 distinct spatial patterns in the optimal fitted Gaussian parameters (Fig. 2). Some of these patterns
471 were similar for the different parameters. In particular, there were zones of generally more intense
472 fires (e.g., South Sudan, northern Central African Republic, Botswana, Namibia and parts of Angola
473 and the Democratic Republic of Congo (DRC)), showing relatively high values of ρ_{peak} , ρ_{base} and σ
474 compared to other zones where values for all three parameters were relatively low (e.g., Zambia,
475 Mozambique, Tanzania, Nigeria and Cameroon). On top of this general pattern, a clear gradient is
476 visible as you move from drier to more humid regions, seen most clearly when moving from Namibia
477 via Angola to DRC. In more humid savannas, when fuel conditions were optimal, high ρ_{peak} values
478 could be reached but fire duration over the day was generally short and night time FRP values were
479 more likely to fall below the SEVIRI FRP detection threshold (Fig. 2). h_{peak} varied considerably over the
480 study region, with areas showing most fire activity late in the afternoon generally in more humid or
481 forested regions but also in some more arid regions (Fig. 2d).

482
483 Table 1 shows the land cover-averaged values and SD of the results presented in Fig. 2. In addition
484 we calculated the ratio of the mean SEVIRI FRP at MODIS daytime detection opportunities to the
485 maximum daytime FRP ρ_{peak} . These results were used in the climatological modelling approach that
486 combined the fire diurnal cycle climatology with observations made at the MODIS sampling times to
487 derive the daily fire diurnal cycle predictions (Sect. 3.5). More intense fires with long duration and
488 high peak values were associated with fires in shrublands, savannas and grasslands, while a more
489 pronounced fire diurnal cycle was present in more humid woody savannas or tropical forests. For σ ,
490 ρ_{peak} and ρ_{base} SD was typically about half of the average value, while SD of h_{peak} was largest for
491 temperate forests, shrublands and grasslands. The ratio of mean daytime FRP made at the MODIS
492 sampling times and ρ_{peak} was relatively constant for various land cover types with ρ_{peak} generally
493 about three times as large as the mean FRP at the daytime MODIS detection opportunities (Table 1).

494
495 In order to better understand the spatial distribution of the fire diurnal cycle features, we studied
496 characteristics of the fire regime that were expected to be related to fuel properties and the diurnal
497 cycle (Fig. 3a, c and d). To guide the interpretation we have included a land cover map, partly
498 governing fuel loads, in Fig. 3b. Annual emitted FRE varied widely over the study region, and highest
499 values were found in the savannas and woody savannas (compare Fig. 3a with b) and coincided with
500 regions of large fire size and/or a high number of annual fire days (compare Fig. 3a with c and d).
501 Similarities with characteristics of the fire diurnal cycle were also found, the earlier mentioned zones
502 of generally more intense fires (high values of ρ_{peak} , ρ_{base} and σ) often coincided with regions of large
503 fire size (Figs. 2a–c and 3c). In the more humid tropical areas, high ρ_{peak} values occurred in areas of
504 relatively large fire size and/or a high number of annual fire days (Figs. 2a and 3c, d).

505
506 The relative fraction of FRE emitted on days that SEVIRI data sampled at MODIS observation times
507 did not observe active fires is an important factor affecting model performance, and showed similar
508 spatial patterns as σ , indicating that duration of fires over the day plays an important role (Figs. 2c
509 and 4a). In addition, the geographical location and cloud cover during the burning season played a
510 role by affecting the effective number of daily MODIS observations (Fig. 4b). The peak hour of fire
511 activity also played a role, and especially in more humid areas with frequent cloud cover and late
512 afternoon fire activity sometimes over 50% of FRE was emitted on days without any SEVIRI active fire
513 detections at MODIS detection opportunities (compare Figs. 2d and 4a). The most important biomass
514 burning regions were typically characterized by relatively long fire duration over the day (Fig. 2c) and
515 the effect of omission of active fires on continental scale FRE estimates was therefore relatively low
516 (cf. Fig. 3a, 4a and 5). However, frequent omission of relatively small fires of short duration may
517 strongly affect FRE estimates for some regions (Fig. 5). These results clearly demonstrate the value of
518 the data provided by the very high temporal resolution geostationary systems, even though they are
519 unable to resolve and detect fire pixels as low in FRP as those from polar orbiters (Roberts and
520 Wooster, 2008).

521

522 **4.2 Model evaluation**

523

524 To evaluate the two modelling approaches that estimated FRE from SEVIRI data only at the MODIS
525 sampling times we started with comparing the spatial distribution of mean estimated FRE for each
526 method with the cloud corrected SEVIRI FRE calculated using the entire hourly, 0.1° SEVIRI FRP
527 dataset (Fig. 5). The persistent approach resulted in a general overestimation of FRE, while the
528 climatological approach showed overall good performance in terms of total estimated FRE when
529 compared to the full SEVIRI dataset. Moreover, the more narrow distribution of modelled FRE as a
530 fraction of SEVIRI FRE by the climatological approach as opposed to the persistent approach suggests
531 that results are not only more accurate but also more precise (Fig. 5). While this reflects the general
532 pattern, the performance bias was not homogeneous over the region. The persistent approach
533 showed best results for regions with long daytime fire durations (i.e., large σ) and with a late peak in
534 fire activity; and although performing generally better, the climatological approach showed a general
535 underestimation for areas of relatively late peak fire activity (compare Figs. 2 and 5). To a certain
536 extent these regional differences correspond to the distribution of the different land cover types
537 (Table 2). For example, for temperate forests and shrublands the persistent modelling approach
538 showed notably better comparison to the FRE derived via the entire SEVIRI dataset, while the
539 climatological modelling approach overestimated FRE.

540

541 Equally important as the absolute FRE estimates shown in Fig. 5 and Table 2 are their temporal
542 dynamics. Figure 6 shows regional daily budgets for several study regions with different geographical
543 positions and land cover. Similar to the results in Fig. 5, we found a general overestimation by the
544 persistent approach, and better overall estimation by the climatological approach. Overestimation of
545 the persistent approach was occurring mostly in the tropics (e.g., Nigeria and DRC), where also
546 stronger day to day variability was observed as compared to that derived with the complete SEVIRI
547 data or the other modelling approaches (Fig. 5). The climatological approach showed a small delay in
548 FRE estimations compared to the complete SEVIRI dataset.

549

550 To further test the ability of the two modelling approaches to allocate FRE to the individual grid cells
551 at the right moment in time, correlation coefficients were calculated. Table 3 shows Pearson's r
552 between SEVIRI and the two modelling approaches at four spatiotemporal resolutions (0.1° and 1°
553 spatial and hourly and daily temporal resolution). A striking increase in correlation was observed
554 when aggregating model results both temporally or spatially. Freeborn et al. (2009, 2011) previously
555 demonstrated the value of such spatial aggregation when deriving relationships between SEVIRI and
556 MODIS datasets, and this technique is currently used within the near real-time SEVIRI FRP-GRID
557 products produced by the LSA SAF from the SEVIRI FRP-PIXEL data (Wooster et al., 2015). At 0.1°
558 resolution the best correlations were found for shrublands and savannas while for aggregated data
559 best performance was found for woody savannas and savannas. At hourly resolution, the
560 climatological approach generally performed better than the persistent approach. However, at 0.1°
561 daily the persistent approach performed best while at 1° spatial resolution the persistent and
562 climatological approaches did equally well.

563

564 **5 Discussion**

565

566 Unlike biomass burning emission inventories based on burned area, inventories using active fire
567 observations from Earth Observation satellites can be produced in near real time (Freitas et al., 2005;
568 Reid et al., 2009; Sofiev et al., 2009; Wiedinmyer et al., 2011; Kaiser et al., 2012; Darmenov and da
569 Silva, 2013). The near real time emissions inventories are, at present, generally based on active fire
570 data from the MODIS instruments operating onboard the Terra and Aqua polar orbiting satellites.
571 The FRP observations of MODIS are almost without saturation, operating day and night, with a
572 reasonable spatial resolution and with new observations available for any location at least a few
573 times every day – cloud cover permitting. However, it is well known that fire activity in most regions
574 follows a clear daily cycle (e.g., Roberts et al., 2009; Vermote et al., 2009). Consequently, the FRP
575 measures derived from intermittent polar orbiting MODIS observations are often not fully and
576 directly representative of the actually daily fire activity (Fig. 1; Giglio, 2007; Vermote et al., 2009;
577 Freeborn et al., 2011). Although several approaches have been developed to obtain more accurate
578 estimations of FRE from the limited temporal sampling of FRP provided by MODIS (e.g., Ellicott et al.,
579 2009; Freeborn et al., 2009, 2011; Vermote et al., 2009), they are all best suited to be used with
580 previously collected and/or aggregated FRP data, and none can be readily implemented at high
581 spatiotemporal resolution in near real time. For this reason, most current global emission inventories
582 produced in near real time actually ignore fire diurnal dynamics completely (e.g., Kaiser et al., 2012),
583 and this results in large biases in the FRE budgets (Ellicott et al., 2009; Zhang et al., 2012).

584

585 Here we start discussing the spatial distribution of the fire diurnal cycle, and its drivers (Sect. 5.1).
586 Building on previous work, we explored two new methods to estimate hourly FRE in near real time
587 from observations made by SEVIRI at MODIS detection opportunities. The results illustrate how
588 MODIS observations might be used to calculate hourly FRE, and where errors can be expected due to
589 the diurnal cycle and the limited temporal sampling provided by MODIS (Sect. 5.2).

590

591 **5.1 Exploring the fire diurnal cycle using a Gaussian function**

592

593 The fire diurnal cycle characteristics were explored by fitting of a Gaussian function to the hourly
594 SEVIRI time-series. Vermote et al. (2009) and Ellicott et al. (2009) found that at a 0.5° monthly
595 resolution the fire diurnal cycle can be described by a Gaussian function, using MODIS observations
596 to resolve the unknown parameters. They choose the spatiotemporal size of the study regions such
597 that a statistical representative number of fires and MODIS FRP detections were included, and the
598 observations covered the full range of MODIS view angles – since the sensitivity of MODIS to fire
599 depends upon this (Vermote et al., 2009). Although later work showed that in fact fire activity may
600 be somewhat skewed in the afternoon, here we found that even at a high spatiotemporal resolution
601 (0.1°; hourly) a Gaussian function provides a fairly robust description of the fire diurnal cycle.
602 However, at 0.1° hourly resolution, SEVIRI data sampled at the MODIS detection opportunities does
603 not always provide enough information to adequately depict fire activity for an individual grid cell
604 and day (Fig. 1). Moreover, the spatiotemporal scale at which we observe the fire diurnal cycle has a
605 significant impact on its shape. When moving to a coarser spatiotemporal resolution, the shape of
606 the diurnal cycle likely becomes wider, with less pronounced peaks. This is mostly a consequence of
607 the spatiotemporal variation in hour of peak fire activity of the individual fires or fire days (Fig. 1).
608 Therefore, typical values of the parameters of the Gaussian found in this study (Fig. 2) do not
609 necessarily correspond to typical values found by earlier studies (e.g., Roberts et al., 2009; Vermote
610 et al., 2009), who used much larger sample sizes (i.e., spatiotemporal resolutions). Likewise the
611 results presented here are not necessarily representative for individual fires.

612

613 Although the shape of the “average” fire diurnal cycle is scale dependent, regional patterns in the
614 diurnal cycle characteristics (Fig. 2) remain similar over different scales, and therefore we found
615 similar land cover dependent characteristics as previous studies. For example, shrublands and
616 grasslands generally faced drier conditions when burning than did woody savannas or tropical forest,
617 and therefore fire activity typically continued longer over the day and the hour of peak fire activity
618 was generally located later in the afternoon (Fig. 2; Table 1; Giglio, 2007; Roberts et al., 2009). For
619 the same reason, temperate and boreal forests have been reported to show a more pronounced
620 diurnal cycle than grasslands (Fig. 2; Sofiev et al., 2013; Konovalov et al., 2014). Building on the land
621 cover based analysis of Roberts et al. (2009), we provide a first analysis of the spatial distribution of
622 the fire diurnal cycle.

623

624 The three parameters determining the shape of the Gaussian can be used to visualize the spatial
625 distribution of the fire diurnal cycle. The daily FRP-maximum is given by ρ_{peak} , fire duration over the
626 day by σ , and the baseline FRP by ρ_{base} . Similar spatial patterns were found for all three parameters
627 mentioned above (Fig. 2a, b and c). This indicates that there are zones of generally more “intense”
628 fires with high ρ_{peak} , large σ and higher ρ_{base} , while other zones are characterised by lower intensity
629 fires. In land cover classes where most of the fires were grass fuelled (grasslands, savannas and
630 woody savannas), a considerable part of the spatial variation in fire diurnal cycle could be explained
631 by fire size (see Sect. 2.4; Figs. 2 and 3). Large fires were often found in frequently burnt and/or more
632 arid areas (Fig. 3a) where high fuel connectivity, low fuel density and low fuel moisture allow
633 relatively fast moving fires with large fire fronts to form (Hély et al., 2003; Sow et al., 2013). Besides
634 fire size and land cover, part of the variability in the fire diurnal cycle could be explained by a
635 gradient in diurnal weather conditions. Grass fuelled large fires were also common in the more

636 humid savannas of southern Africa, but here nighttime weather conditions appear to become rather
637 unfavourable for fire (Figs. 2b and 3c). In humid savannas ρ_{peak} values were not solely associated with
638 large fire size, but also with areas showing a high number of annual days with fire activity and may be
639 explained by several relatively small fires burning at the time. The high number of fire days may
640 indicate a larger number of fire ignitions and/or that fires are spreading at a slower rate due to the
641 more pronounced fire diurnal cycle, higher humidity, or higher fuel density (Hély et al., 2003; Sow et
642 al., 2013). Finally, in the Mediterranean basin the relatively low fire return period, and consequently
643 higher fuel density, may also cause relative intense fires with long duration over the day (Fig. 2;
644 Archibald et al., 2013).

645
646 The peak hour of fire activity found here corresponds to the moment of day at which 50% of the total
647 FRE has been emitted (assuming $\rho_{base} \ll \rho_{peak}$), and it did not always correspond to the peak hour
648 of fire activity found by previous studies (Fig. 2d; e.g., Giglio, 2007; Roberts et al., 2009; Vermote et
649 al., 2009). In general most FRE was emitted during the afternoon, and clear spatial patterns were
650 present in the typical peak hour of the Gaussian. High values of h_{peak} were found in regions of higher
651 fuel density or in more arid areas where fires could spread over large areas (Figs. 2d and 3). In arid
652 regions with large typical fire sizes, fire spread was often fast and a 0.1° grid cell only corresponded
653 to a part of the actual fire resulting in large variation in h_{peak} between neighbouring grid cells (Fig. 2d
654 and Table 1).
655

656 **5.2 Model performance and the MODIS sampling design**

657
658 Data assimilation and two modelling approaches, were used to estimate hourly FRE from SEVIRI FRP
659 data sampled at the times of MODIS detection opportunities. Here we start discussing the
660 performance of the different methods with respect to their total FRE estimates and daily regional
661 FRE estimations. Then we discuss the more uncertain model performance at higher spatiotemporal
662 resolutions.

663
664 The persistent approach is comparable to a direct hourly extension of the current GFAS methods
665 (Kaiser et al., 2012), where the fire diurnal cycle is ignored and the predicted FRP for each hour is
666 equal to that of the last FRP observation. This led to a general overestimation of daily FRE because
667 the 13:30 LT temporal sampling time of MODIS is relatively close to the peak hour of daily fire
668 activity, and therefore not very representative of the full period until the next observation at 22:30
669 LT (Figs. 2d and 5; Table 2). Moving away from the equator, the number of daily MODIS observations
670 increases due to orbital convergence at higher latitudes, and consequently the model performance
671 improved (Figs. 4b, 5 and 6; Giglio et al., 2006; Reid et al., 2009). Additional inclusion of daytime
672 observations due to orbital convergence will typically be somewhat earlier or later in the afternoon
673 and may therefore lower the FRE estimation. In the persistent approach, missing nighttime
674 observations may cause an overestimation and missing daytime observation an underestimation of
675 daily FRE, resulting in erroneous regional day-to-day variations in FRE estimates in the tropics (Fig. 6).
676 Following previous research, we found that due to the spatiotemporal variation of the fire diurnal
677 cycle FRE was overestimated more for some land cover types than for others (Table 2; Freeborn et
678 al., 2011). Land cover classes that typically showed longer fire durations (Fig. 2c) with peak fire
679 activity later in the afternoon (Fig. 2d) were not as much overestimated as land cover classes with

680 more pronounced fire diurnal cycles (Figs. 5 and 6; Table 2). However, part of this effect likely stems
681 from these land covers mostly being located in the more frequently observed higher latitudes of our
682 study region. Although the persistent method is not directly comparable to the methods of widely
683 used emission inventories like GFAS or QFED (Kaiser et al., 2012; Darmenov and da Silva, 2013), they
684 likely introduce similar errors by ignoring the fire diurnal cycle.

685
686 The climatological approach showed better performance in terms of absolute FRE estimations, while
687 also better able to reproduce its spatial variability (Fig. 5). In contrast to the persistent approach, the
688 hourly predictions were based on the last 24h of fire activity, enabling more realistic gap filling during
689 periods without observations. This resulted in an advantage during periods of cloud cover or missing
690 observations due to the satellite orbits, but because of the low number of actual daily observations
691 the climatological approach had the tendency to continue predicting fire activity after fires had
692 ceased, seen as a small delay in the signals in Fig. 6.

693
694 An additional criterion to evaluate the model performance was the correlation between the
695 modelling approaches and the SEVIRI data at different spatiotemporal scales. Correlation between
696 the modelled and SEVIRI time-series improved considerably when moving from hourly to daily
697 resolution, showing that the models were better able to estimate daily budgets than the distribution
698 of fire activity over the day. These differences may be explained by the inability of the models to
699 correctly estimate the hour of peak fire activity, a fire diurnal cycle that is not well represented by a
700 Gaussian function, or in the case of small fires the fire diurnal cycle may not be fully detected by the
701 SEVIRI instrument. Because of the large day-to-day variation in the fire diurnal cycle and the FRP
702 measurements limited to the time of the MODIS overpasses, the individual FRP observations have a
703 low precision (i.e., large random error) and omission (i.e., non detection) of fires is frequent (Figs. 1
704 and 4), resulting in low correlation at high spatiotemporal scales (Table 3). Because fires rarely occur
705 on their own and generally form part of a regional pattern (Bella et al., 2006), the correlation
706 increased considerably when accumulating results to a 1° spatial scale. For the same reason model
707 performance was found to be best in savannas and woody savannas, where the highest number of
708 fires occur and the sample size is thus largest, or in areas of large fire size where omission was
709 relatively low. Model performance was therefore best when optimal burning conditions were
710 reached, often coinciding with the peak of the burning season. Because often only a reasonably large
711 sample of observations made at the MODIS detection opportunities is actually representative of fire
712 activity in a certain region, the added value of the 0.1° spatial resolution (e.g., GFASv1.1/1.2) is
713 somewhat limited compared to a coarser 0.5° spatial resolution (e.g., GFASv1.0).

714
715 Overall, using the climatological approach resulted in the best model performance, although in
716 specific cases using the persistent approach showed better results. For example, at 0.1° spatial and
717 daily temporal resolution the persistent approach performed best, likely because it only predicts fire
718 activity on days of actual fires while the climatological approach may predict fire activity with some
719 delay. Also the climatological approach used mean values for the fire duration and may therefore
720 overestimate FRE from smaller fires while underestimating the larger fires. Despite the improved
721 results of the climatological approach as opposed to the persistent approach, estimating FRE in near
722 real time based on MODIS observations remains challenging, especially at high spatiotemporal
723 resolutions. Largest uncertainties originate from the high spatiotemporal variability of the fire diurnal
724 cycle combined with the limited number of daily MODIS detection opportunities. Moreover, the fire

725 diurnal cycle as analyzed here may to some extent be affected by the inability of SEVIRI to detect the
726 smallest fires, along with other sources of uncertainty in the FRP observations (Wooster et al., 2015;
727 Roberts et al., 2015). Finally, the characterization of the fire diurnal cycle and discussion of its
728 spatiotemporal drivers presented here provide a first step to upscale the climatological model to a
729 global scale, but a better understanding of the fire diurnal cycle and its drivers for other regions of
730 the globe remains an important issue.

731
732 Within GFAS, to handle the uncertainties introduced into the MODIS-derived FRE estimates by
733 neglecting the diurnal cycle influence, the estimated FRE is converted into estimates of dry matter
734 burned (DM) using land cover-specific conversion factors. These were derived via comparison of
735 long-term monthly FRE estimates to the DM estimates calculated over the same period by the Global
736 Fire Emissions Database (GFED 3.1; van der Werf et al., 2010; Kaiser et al., 2012). It is currently
737 assumed that by allowing the conversion factors to vary with land cover type the impact of any land
738 cover-varying diurnal cycle is also incorporated, reducing the influence of the diurnal cycle. The
739 issues discussed above, along with the accuracy of the GFED DM calculations, which are for example
740 affected by the quality of the burned area product and the biochemical models used, all influence
741 values of the land cover-specific FRE-to-DM conversions factors presented by Kaiser et al. (2012).

742
743 Wooster et al. (2005) and Freeborn et al. (2008) previously explored the conversion factors between
744 FRE and DM using small scale experiments, and found that they appeared relatively independent of
745 vegetation type. However, when moving to the satellite-scale there are additional factors influencing
746 this FRE-to-DM relationship, for example the fire regime of an area and the degree to which MODIS
747 misses the lowest FRP fires, and the canopy density of trees that might obscure some of the thermal
748 radiation being emitted by fires burning in the ground fuels (Freeborn et al., 2014). The thermal
749 radiation recorded in satellite products is additionally reduced by cloud cover and erroneous flagging
750 of smoke as clouds during data processing. Konovalov et al. (2014) nevertheless found FRE-to-DM
751 relationships relatively similar to those of the earlier small-scale experiments when using
752 atmospheric observations and biomass burning trace gas and aerosol emissions factors to estimate
753 fuel consumption. Exploring methods to incorporate the fire diurnal cycle in the GFAS global FRP-
754 based near real time emission inventory is a first step in taking into account some of these issues in
755 order to improve global FRE estimates made at relatively high spatiotemporal resolutions, and
756 hopefully also in reconciling some of the differences in current emission inventories.

757

758 **6 Conclusions**

759
760 Emission inventories based on FRP observations have great potential to improve biomass burning
761 emission estimates, by eliminating the need for modelling of fuel loads and fuel consumption, and
762 can be produced in near real time. However, to date uncertainties in FRE estimation remain high
763 when using polar orbiting FRP datasets, largely due to difficulties in combining the limited temporal
764 resolution observations and knowledge about the fire diurnal cycle. Geostationary data can alleviate
765 this issue, but brings its own problems related to the non-detection of the lower FRP fires due to the
766 coarse spatial resolution of the geostationary observations. Geostationary dataset are also not global
767 in extent. Here we explored the spatial dependencies of the fire diurnal cycle and its impact on active
768 fire detections made at the time of MODIS overpasses. Two modelling approaches were developed

769 to derive hourly FRE estimates based on data-assimilation and SEVIRI FRP observations subsampled
770 at MODIS detection opportunities. The first approach ignored the fire diurnal cycle assuming
771 persistent fire activity between two MODIS detection opportunities, while the second approach
772 combined prior knowledge of the fire diurnal cycle with active fire observations at MODIS detection
773 opportunities to simulate the fire diurnal cycle. Both approaches were evaluated against the actual
774 hourly FRP observations made by SEVIRI. Our main conclusions are:

- 775 1. We considered various drivers of the spatial distribution of fire diurnal cycle: dominant land
776 cover, fire size, annual number of fire days, and diurnal climate conditions and found that all
777 played a role. The strong relation between fire size and fire diurnal cycle for grass fuelled
778 fires, and the climatic gradient in diurnal cycle, indicate that using fuel characteristics rather
779 than land cover alone to characterize the fire diurnal cycle provides a potential pathway to
780 improve these estimates. Here we showed that this information can partly be obtained by
781 studying the fire characteristics, such as fire size, which are contained within the remote
782 sensing data themselves.
- 783 2. Ignoring the fire diurnal cycle may cause structural errors in FRE estimates, and likely results
784 in a general overestimation of FRE due to the timing of the MODIS overpasses. The errors
785 vary regionally, mostly due to variations in the fire diurnal cycle, while results get more
786 accurate at higher latitudes due to the increasing number of daily MODIS detection
787 opportunities caused by orbital convergence.
- 788 3. Due to the large day-to-day variations in the fire diurnal cycle at the grid cell level, and the
789 scarce number of MODIS observations of any one location per day, daily FRP fields calculated
790 from observations made at MODIS detection opportunities are characterized by low
791 precision (i.e., observations are not representative for daily fire activity) and high omission
792 (i.e., non observation of fires). Therefore a sufficiently large sample size of MODIS
793 observations is required to accurately estimate FRE, as shown earlier by Freeborn et al.
794 (2011). In zones of frequent fires, where fires are generally part of a regional biomass
795 burning pattern, model performance greatly improved when moving to a coarser scale,
796 increasing the sample size. Model performance was also considerably better for zones of
797 relatively large fires that were characterized by low omission. Production of emission
798 inventories at very high spatiotemporal resolution using data from a limited number of low-
799 Earth orbit satellite observations may therefore provide somewhat restricted added value
800 compared to those derived at coarser spatiotemporal scales.
- 801 4. Relative overrepresentation of day- or nighttime FRP observations may cause large day to
802 day variations in estimated FRE when the diurnal cycle is ignored.
- 803 5. The way we observe the fire diurnal cycle is scale dependent, mostly because of the large
804 variation in fire diurnal cycle, even within the same grid cell between different days.

805
806 We recommend implementing the climatological model within GFAS in Copernicus Atmosphere
807 Services in order to improve global and regional FRE estimates and further reconcile emission
808 estimates from the various different inventories currently available.

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814 283576 and 633080).

815

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1020 Table 1. Mean values of the parameters of the Gaussian function per land cover type (excluding days
 1021 of cloud cover and weighted by FRE), SD are shown in parenthesis. Values of σ and the ratio of ρ_{peak}
 1022 and mean day-time FRP at MODIS detection opportunities ($MODIS_{mean}$) were used within the
 1023 climatological approach to model hourly FRP (see Sect. 3.5).

Land cover	σ (hour)	ρ_{peak} (MW)	ρ_{base} (MW)	h_{peak} (hour)	$\rho_{peak}/MODIS_{mean}$ (-)
Temperate forest	1.14 (0.55)	846 (392)	24.2 (12.7)	13.31 (4.50)	3.17
Tropical forest	0.85 (0.45)	1364 (863)	27.3 (19.6)	13.34 (2.57)	3.03
Woody savanna	0.94 (0.50)	1501 (934)	21.1 (16.8)	13.21 (2.08)	3.07
Savanna	1.09 (0.53)	1711 (899)	39.0 (25.5)	13.08 (2.58)	2.88
Shrubland	1.35 (0.63)	3079 (1552)	108.9 (56.9)	13.16 (4.46)	2.87
Grassland	1.06 (0.53)	1642 (863)	37.3 (21.1)	12.95 (4.44)	3.08
Cropland	0.95 (0.48)	1259 (705)	23.9 (16.0)	13.33 (3.22)	2.94

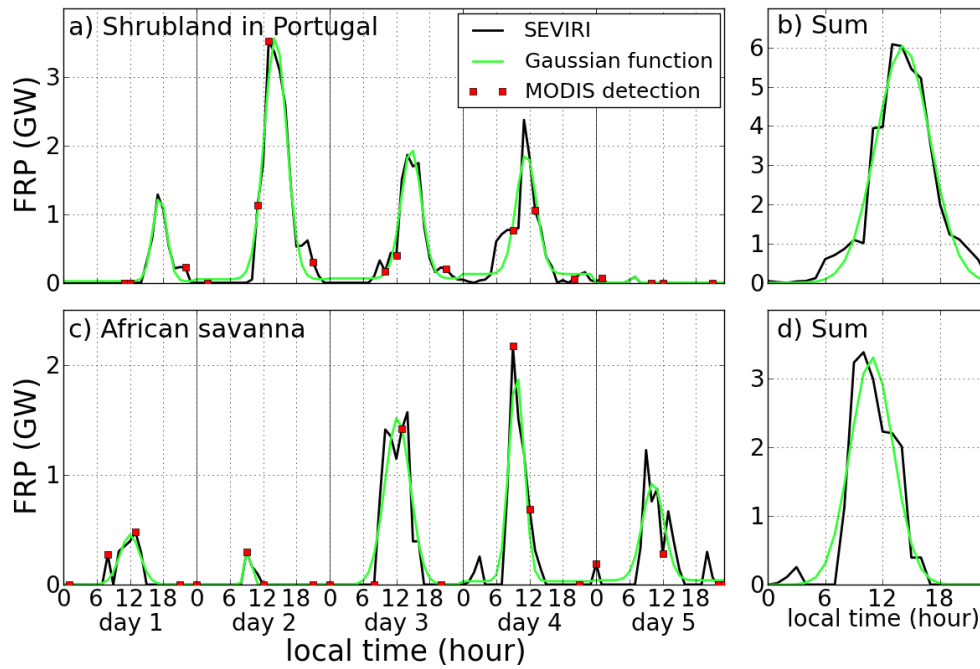
1024 Table 2. Estimated annual FRE during 2010–2012 by the two model approaches as percentage of
 1025 SEVIRI FRE (cloud corrected).
 1026

Land cover	SEVIRI (PJ yr ⁻¹)	Persistent (%)	Climatological (%)
Temperate forest	2.9	98	118
Tropical forest	61.3	179	98
Woody savanna	1513.2	174	93
Savanna	990.7	155	99
Shrubland	91.7	120	115
Grassland	106.5	125	108
Cropland	74.5	147	90
Total	2841.9	163	97

1027 Table 3. Pearson's r between hourly and daily FRE as observed by SEVIRI and estimated by the two
 1028 modelling approaches. Correlation is calculated for two spatial scales, the original 0.1° resolution and
 1029 a 1° aggregated resolution (in parentheses) to test regional model performance.
 1030

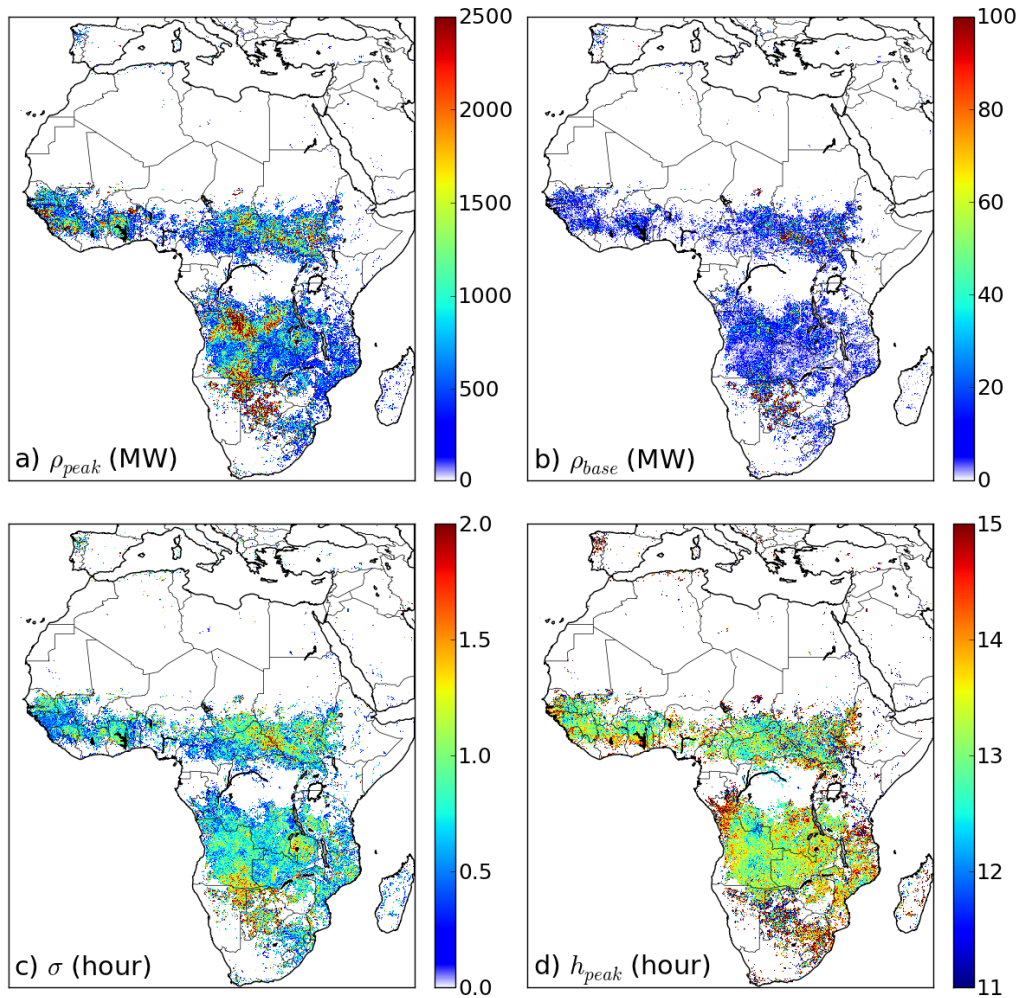
Land cover	Persistent hourly	Climatological hourly	Persistent daily	Climatological daily
Temperate forest	0.24 (0.33)	0.20 (0.32)	0.44 (0.50)	0.21 (0.39)
Tropical forest	0.13 (0.25)	0.15 (0.27)	0.32 (0.41)	0.16 (0.41)
Woody savanna	0.19 (0.44)	0.20 (0.52)	0.48 (0.80)	0.25 (0.79)
Savanna	0.25 (0.45)	0.25 (0.51)	0.54 (0.78)	0.30 (0.76)
Shrubland	0.35 (0.47)	0.32 (0.47)	0.61 (0.63)	0.37 (0.60)
Grassland	0.22 (0.32)	0.20 (0.35)	0.46 (0.55)	0.22 (0.52)
Cropland	0.19 (0.32)	0.17 (0.36)	0.42 (0.61)	0.18 (0.60)
Total	0.22 (0.43)	0.22 (0.50)	0.50 (0.76)	0.27 (0.75)

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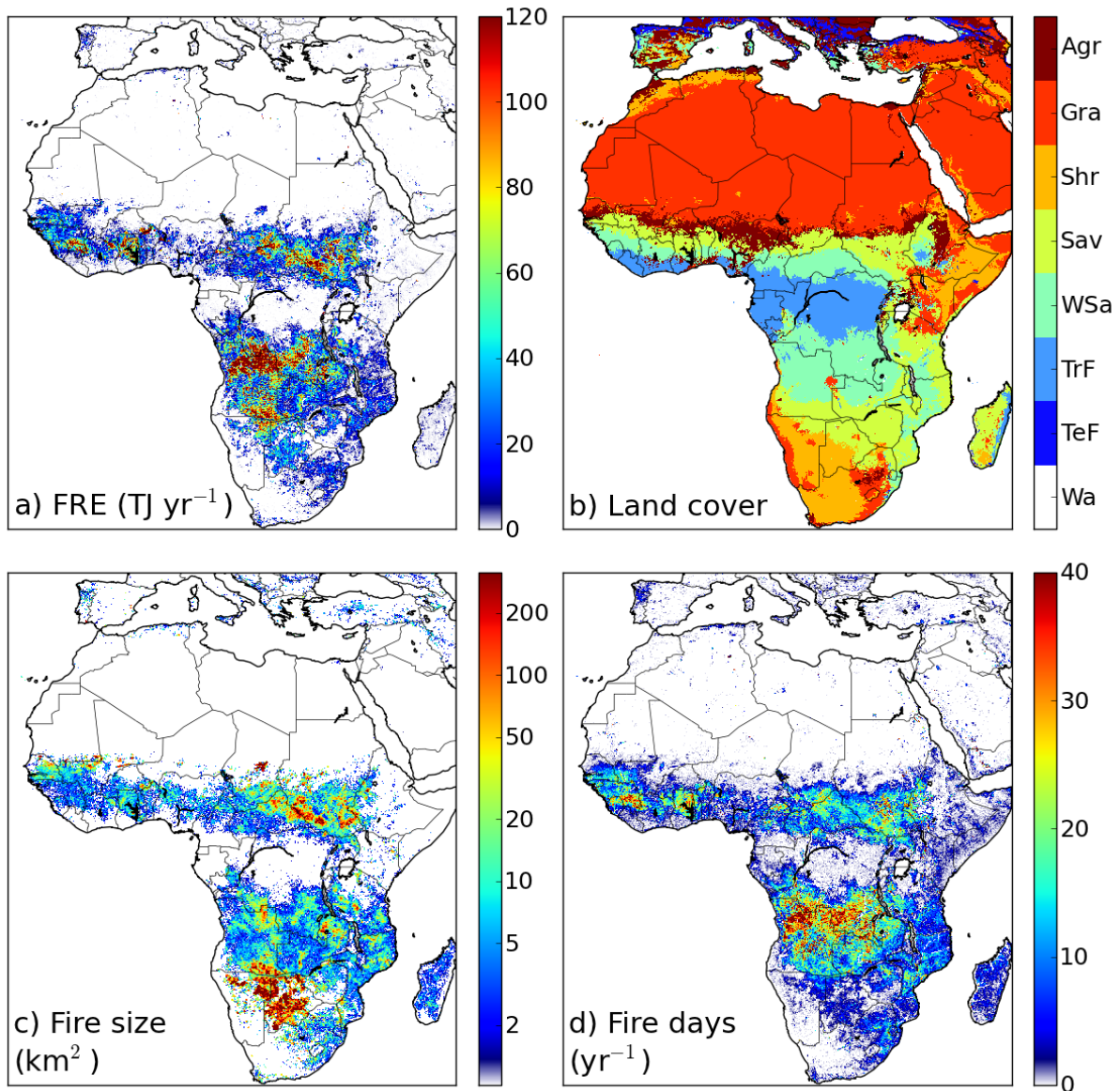
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Figure 1. Hourly-mean FRP time-series derived from SEVIRI data, the same data but only sampled at MODIS detection opportunities, and an optimally fitted Gaussian function fitted to the full SEVIRI FRP time-series. These two examples are for a 0.1° shrubland grid cell in Portugal (a, b) and a 0.1° savanna grid cell in Africa (c, d). (a, c) represent the hourly time-series and (b, d) the aggregated fire diurnal cycle over the 5 study days. Time is indicated as local time.

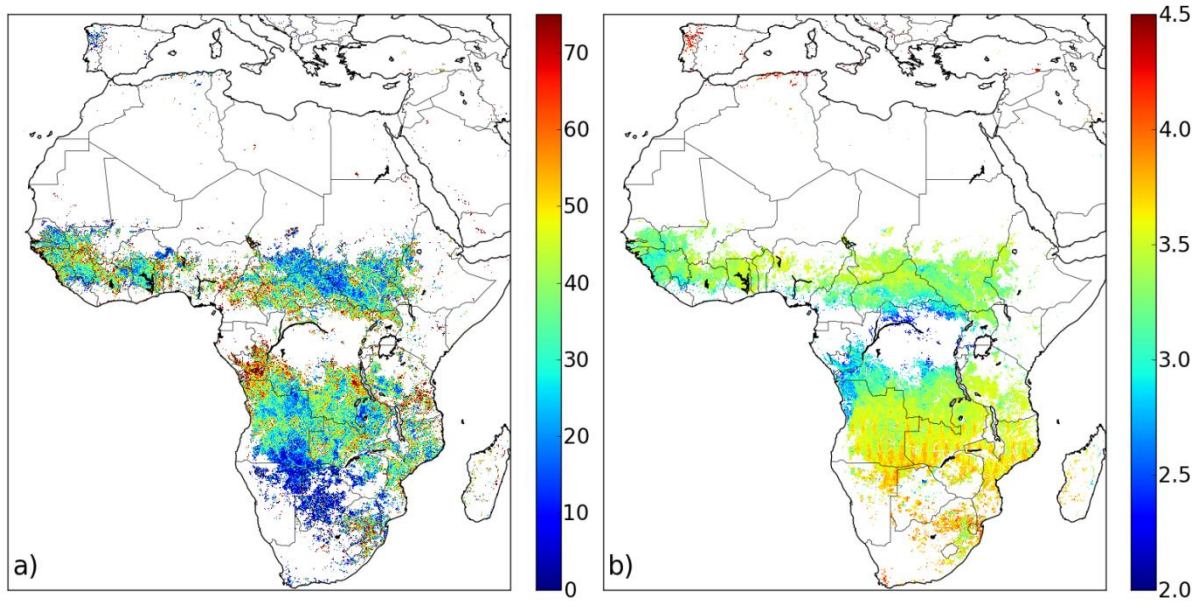


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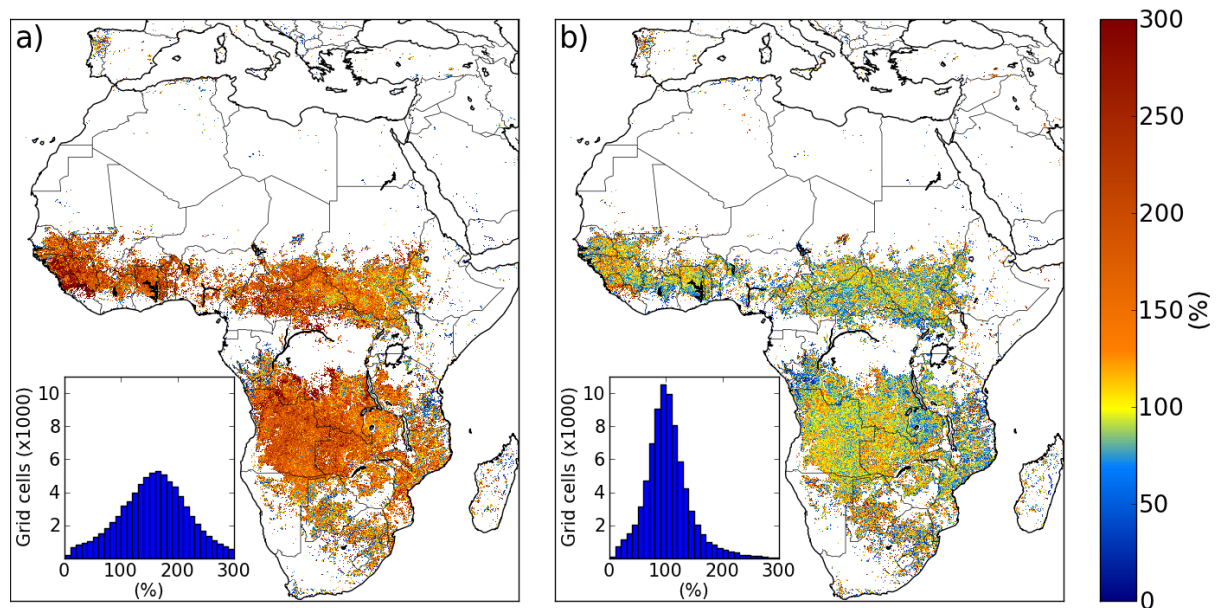
Figure 2. Weighted mean values of parameters of the optimally fitted Gaussian model for each 0.1° grid cell, including all cloud free days during the study period. (a) Peak daytime FRP ρ_{peak} , (b) night time FRP ρ_{base} , (c) SD of the FRE distribution over the day σ (related to the fire duration over the day, or width of the diurnal cycle), and (d) hour of peak fire activity h_{peak} (local time). Grid cells with emitted energy below 5 MJ over the study period (approximately the FRE emitted during one small fire event) were excluded from the figure to facilitate interpretation.



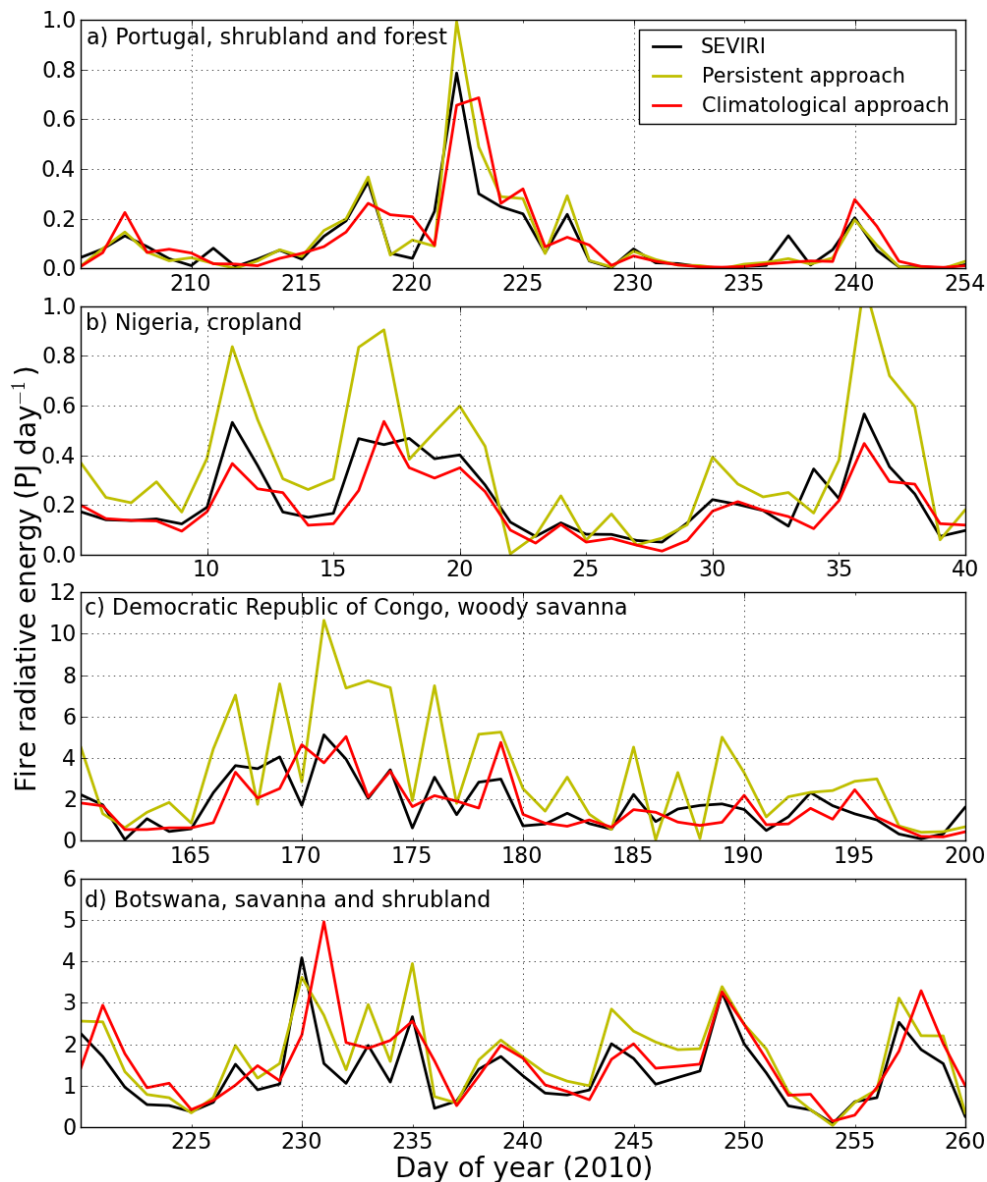
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 1048 Figure 3. Characteristics of the fire regime and fuel types based on 2010 – 2012 data. (a) Mean
 1049 annual FRE per 0.1° grid cell, (b) dominant land cover type, (c) fire size (i.e., weighted mean burned
 1050 area per fire event) and (d) mean annual number of days with fire activity per grid cell over the study
 1051 period. Abbreviations of land cover classes: water (Wa), temperate forest (TeF), tropical forest (TrF),
 1052 woody savanna (WSa), savanna (Sav), shrubland (Shr), grassland (Gra) and agriculture (Agr).
 1053



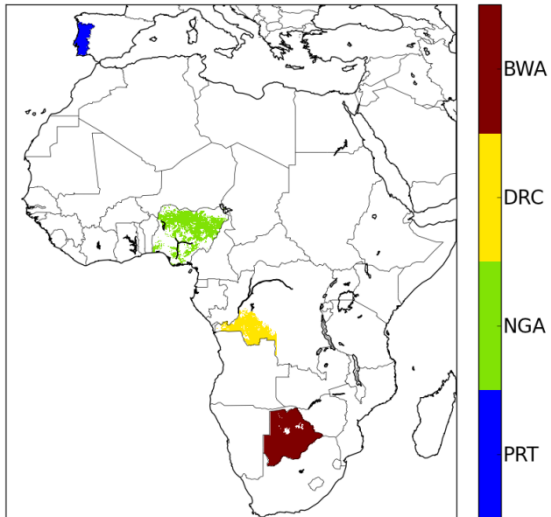
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 1055 Figure 4. Detection of fire activity at MODIS detection opportunities. (a) Percentage of FRE emitted
 1056 on days that the SEVIRI instrument did not observe active fires at MODIS overpasses. (b) Number of
 1057 MODIS detection opportunities per day during the burning season (mean over the study period,
 1058 weighted for monthly FRP).
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 1061 Figure 5. Total fire radiative energy (FRE) estimated via the two modelling approaches using SEVIRI
 1062 observations taken at only the MODIS detection opportunities, expressed as fraction of the total FRE
 1063 calculated using the entire set of hourly mean, 0.1° SEVIRI FRP observations (cloud cover corrected).
 1064 (a) Persistent approach, and (b) climatological approach. Distribution of the grid cell values is shown
 1065 in the lower left corners.
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 1068 Figure 6. Daily FRE for four study regions (areas of 85000 to 567000 km²) derived from the complete
 1069 SEVIRI dataset (cloud cover corrected) and estimated by the two modelling approaches developed
 1070 here. (a) Daily FRE for Portugal, mostly including shrublands and temperate forests, (b) fires in
 1071 Nigeria burning in croplands, (c) woody savannas in DRC, and (d) shrublands and savannas in
 1072 Botswana. Study regions are shown in Fig. 7, and land cover was determined using the dominant land
 1073 cover classes (Sect. 2.3; Fig. 3b).



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Figure 7. Study regions used in Fig. 6. Abbreviations refer to: Botswana (BWA), the Democratic Republic of Congo (DRC), Nigeria (NGA) and Portugal (PRT).