

Response to Reviewer Comments:

We thank the two Anonymous Reviewers for their thorough comments; we think they have helped improve the content of the manuscript.

In February (a few months after the manuscript was submitted) we discovered an issue with the treatment of pressure in the model. Our model assumed the pressures supplied by the NASA GMAO office were dry pressures when, in fact, they were moist. We have fixed this error and redone many of the simulations.

The figure below shows the impact of this correction. We found the largest impacts at the surface and the impact on the total column to be relatively small (~ 1.2 ppbv; with respect to CH_4 and other biases) because water vapor concentrations drop off rapidly with increasing altitude. Additionally, the bias in the total column is, for the most part, spatially uniform. This means it would have been folded into our bias correction shown in Figure 2 of the manuscript. We also tested to ensure our updated simulation was still conserving mass.

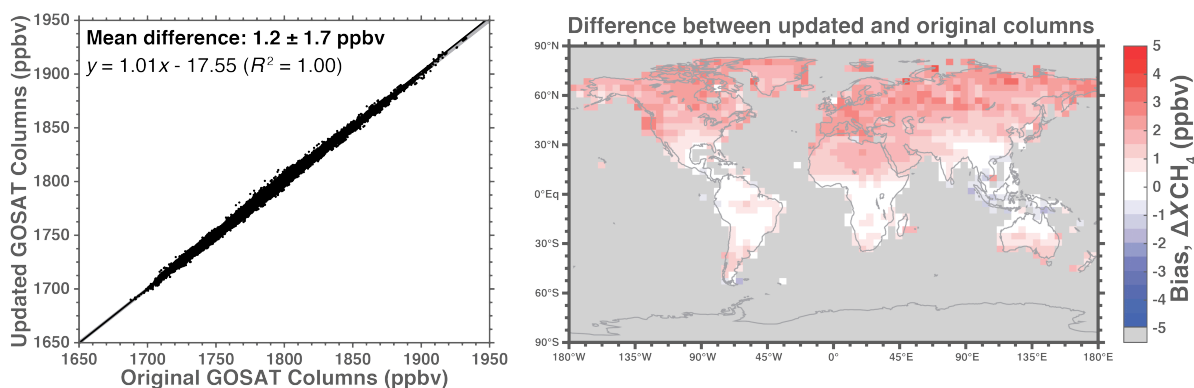


Figure 1: (Left) Scatterplot comparison of the original and updated GOSAT columns. (Right) Spatial differences between the original and updated GOSAT columns.

Furthermore, the TCCON retrievals have been updated to use the GGG2014 version and we have added additional comparisons to the NOAA/ESRL Tall Tower Network.

Reviewer #1 Comments:

1.) Attribution: As currently written, the manuscript implies stronger attribution than actually is possible in the observing-model framework presented. In the abstract and conclusion it is very strongly stated that “We attribute”. In fact, the model-inversion framework does the attribution. Importantly, this relies on the prior distribution of source types, and the assigned uncertainty. As the author’s state later in the manuscript depending on how one constructs the uncertainty either oil/gas or livestock are ‘attributed’ as the largest discrepancy in the US. The author’s should clearly state in the abstract and conclusion that attribution is purely a model product, and relies on accurate spatial distribution of prior sources (or at least distribution of dominant sources) and is dependent on uncertainty assigned by the author’s with expert judgment.

We feel that the attribution from this paper is actually more conservative than past work using similar observation-model frameworks and is a fair attribution for the sources. In the abstract and throughout the manuscript we present a range of attributions (e.g., 29-44% of anthropogenic emissions to livestock, 22-31% to oil/gas, etc.) that is based on inversions with varying specifications of the prior uncertainty. Before presenting the attribution in the paper we explain the underlying assumptions:

Lines 197-200: “We inferred the contributions from different source types to our posterior emissions by assuming that the prior inventory correctly partitions the methane by source type (see Appendix and Table 1) in each 4°×5° grid cell. This does not assume that the global distribution of source types is correct in the prior, only that the local identification of dominant sources is.”

Further, these different specifications of the prior uncertainty provide a possible reconciliation for the different attributions in past work (e.g., Miller et al., 2013 and Wecht et al., 2014) that we discuss at length (Lines 327-336). The main conclusion and an avenue to address this limitation are copied below.

Lines 336-340: “With current prior knowledge it is thus difficult to conclusively attribute the US EPA underestimate to oil/gas or livestock emissions. This limitation could be addressed by a better prior knowledge of the spatial distribution of source types or by use of correlative information (e.g., observations of ethane originating from oil/gas) in the inversion.”

We have changed the text accordingly:

Lines 16-18: “Using prior information on source locations we attribute 29–44% of US anthropogenic methane emissions to livestock, 22–31% to oil/gas, 20% to landfills/waste water, and 11–15% to coal.”

Lines 366-368: “On the basis of prior emission patterns we attribute 22–31% of US anthropogenic methane emissions to oil/gas, 29–44% to livestock, 20% to waste, and 11–15% to coal.”

2.) Representation error: GOSAT has a footprint with a diameter of ~10.5 km. The GEOS-Chem model is either being run at 4x5 or 1/2 by 2/3 degree resolution (~50 x50 km). The author’s never discuss how they address the mismatch between the simulated column enhancements and the observed. Importantly, in regions of strong sources (and topography) and therefore strong XCH₄ gradients, observations at 10.5km can and will see signals averaged out at 50x50km (even worse at 4x5 degrees). How are the authors dealing with this? What are the implications of this? In particular given the non-smooth nature of methane emissions from anthropogenic sources, the resolution of these model runs may not be sufficient to constrain fluxes at the uncertainties reported.

We have attempted to account for the representation error. The observational error variances (which include representation error) are specified based on the residual standard deviations (cf. Heald et al., 2004) unless the retrieval error reported for a given observation is larger than the residual standard deviation, in which case we use the reported observation error. Figure 2b from the manuscript shows a plot of the residual standard deviations for the

global inversion and we discuss this in the updated text:

Lines 165-172: “Observational error variances are estimated following Heald et al. (2004) by using residual standard deviations of the differences between observations and the GEOS-Chem simulations with prior emissions, as shown in Fig. 2b. This method attributes the mean bias on the $4^{\circ}\times 5^{\circ}$ grid to errors in emissions (to be corrected by the inversion) and the residual error to the observational error (including contributions from instrument retrieval, representation, and model transport errors). If the resulting observational error variances are less than the local retrieval error variances reported by Parker et al. (2011) then the latter are used instead. This is the case for 58% of the observations, implying that the observational error is dominated by the instrument retrieval error.”

3.) Transport error: When linking atmospheric observations to fluxes, atmospheric transport is the key integral ingredient. Many regional, continental, and global studies go to great lengths to attempt to quantify the potential role of transport error and its impact on the inverted fluxes. This has not been addressed by the author’s in this manuscript adequately for inverting fluxes of greenhouse gases. GEOS winds are not perfect, and the representation at 4×5 degrees and 50×50 km might greatly impact the analysis and interpretation. What confidence can we have in the transport accuracy? Can we quantify the uncertainty in that term and include it in the inversion? Can we test if a bias error is possible? Many of the cited regional studies over California or North America make efforts to compare different transport as a proxy to understand transport error.

The residual standard deviations will account for some transport error, provided the error is not systematic. We have updated the text accordingly:

Lines 165-169: “...residual standard deviations of the differences between observations and the GEOS-Chem simulations with prior emissions, as shown in Fig. 2b. This method attributes the mean bias on the $4^{\circ}\times 5^{\circ}$ grid to errors in emissions (to be corrected by the inversion) and the residual error to the observational error (including contributions from instrument retrieval, representation, and model transport errors).”

Minor Comments:

1.) Abstract: Line 1-2. The GOSAT observations do not constrain the inversions, but are used in an inversion framework to constrain fluxes please correct wording here.

We have updated the text accordingly:

Lines 1-3: “We use 2009–2011 space-borne methane observations from the Greenhouse Gases Observing SATellite (GOSAT) to estimate global and North American methane emissions with $4^{\circ}\times 5^{\circ}$ and up to $50\text{ km} \times 50\text{ km}$ spatial resolution, respectively.”

2.) Line 3: It is confusing to see degrees and then kilometers for resolution. Since the model is run at $1/2\times 2/3$ at higher resolution, that should be stated here (and ~ 50 km can be in

parentheses, but the \sim is needed since it is only approximate).

We feel that kilometers are more intuitive at that spatial scale. Additionally, we phrased the latter resolution as “**up to** 50 km \times 50 km” because we use the RBFs for the state vector. So we do not need the “ \sim ” because we have the “up to”.

3.) Line 9-10: This makes it appear there is some circularity in the analysis the aircraft/surface data is used to address a bias and is then again used as an independent check. I don't think you can still call this an independent check at this point.

We have updated the text accordingly:

Lines 183-185: “Since we used them previously to justify a bias correction in the comparison between GEOS-Chem and GOSAT, they do not provide a true independent test of the inversion results.”

4.) Page 4497, Line 12: As you cite later, this result has been found and reported previously, so should not be stated in the abstract as if a new, novel finding it is consistent with other studies.

We have updated the text accordingly:

Lines 7-9: “Our global adjoint-based inversion yields a total methane source of 539 Tg a⁻¹ with some important regional corrections to the EDGARv4.2 inventory used as a prior.”

5.) Page 4497, Line 21: The model framework does the attribution (with greater uncertainty that reported)

See our response to the first major comment.

6.) Page 4497, Line 25: “most important anthropogenic greenhouse gas” is more of a subjective statement better to explicitly refer to its climate-relevant role as done in the IPCC report being cited.

We have updated the text accordingly:

Lines 21-22: “Methane (CH₄) emissions have contributed 0.97 W m⁻² in global radiative forcing of climate since pre-industrial times, second only to CO₂ with 1.7 W m⁻² (IPCC, 2013).”

7.) Page 4499, lines 6-14. I am confused a bit here. The large spatial overlap in source types is not a problem you circumvent in fact this also limits you. Furthermore, it is misleading to imply satellite work has only focused on global scales you later cite work where satellite methane observations are analyzed are much smaller scales.

Sources are more likely to overlap at coarse spatial scales (e.g., $4^\circ \times 5^\circ$), by moving to higher resolution some of that overlap can be resolved. We have clarified this in the text.

Lines 57-58: “Spatial overlap is reduced at higher resolution, thus optimizing emissions at high spatial resolution can help improve source attribution.”

As for the latter point, here we are focused on “Previous inversions of methane emissions using satellite data”. We begin by discussing what has generally been done (“mainly focused on coarse scales”) and then discuss work by Wecht et al. (2014a,b) in the latter half of the paragraph that did inversions with satellites at finer spatial scales.

8.) Page 4502, Line 25. This seemingly implicates poor representation of stratospheric CH₄ (and/or tropopause height) rather than a GOSAT artifact could you comment more on this? And what are the implications in regions with significant topography, where there is variance in the tropospheric column contribution to the total column?

We have attempted to investigate this issue using multiple datasets. For example, we have been collaborating with Kat M. Saad (Caltech; Wennberg group) who has been looking at the partial columns with TCCON (Saad et al. 2014) to try and diagnose the source of the bias. There isn’t a conclusive cause for the bias yet. As such, we have not explicitly attributed the bias to GEOS-Chem or GOSAT and have left it open for future investigation. In any case, we have rephrased this discussion to hopefully clarify this section:

Lines 145-147: “This suggests a potential bias in the GEOS-Chem simulation of methane in the polar stratosphere, which warrants further investigation with observations such as TCCON partial columns (Saad et al. 2014; Wang et al. 2014). In any case, we remove the bias using...”

As for the second point, the model uses a terrain-following hybrid pressure-sigma vertical grid. So the model explicitly accounts for topography. We then regrid the model to the satellite vertical grid and convolve the model with the satellite averaging kernel. Additionally, by running the model at higher resolution we are able to more accurately account for variations in topography and don’t need ad hoc topography corrections. Sub-grid scale variations in topography are unlikely to induce a bias because the model uses mean topography. Sub-grid scale variations in topography would lead to increased representation error but that would be captured by the residual SDs (Fig 2b).

9.) Page 4506, line 5. It could be misleading to state the errors are fully characterized. There are a large number of assumptions input including uncertainty levels, lack of systematic errors, and lack of covariances, which are very important in the total error assessment and are not included in the uncertainty range presented.

This simply means we obtain a full posterior distribution instead of the MAP solution. The companion paper (Turner and Jacob, ACPD 2015) discusses this at length. However, we have removed “full” and updated the text accordingly:

Lines 10-11: “...using radial basis functions to achieve high resolution of large sources and provide error characterization.”

Lines 70-71: "...with up to 50km x 50km resolution and error characterization."

10.) Page 4507 line 21-22. I'm not quite sure if state-scale is well-defined constraining California does not imply any other states could have their flux quantified independently CA is very large and is on the ocean so it doesn't have much upwind sources. I would suspect a more accurate statement could be made here about the ability of GOSAT to constrain a specific spatial extent defined by the number of 50x50km boxes that can be constrained together.

We have updated the text accordingly:

Lines 266-267: "By using a longer time record and an optimally defined state vector we achieve much better success."

11.) Page 4508, line16-17. This might not be quite equivalent to assuming the prior distribution is correct but it is close and is highly dependent on the prior distribution.

The reviewer is correct in pointing out that this is dependent on the prior, however the latter part of that sentence explicitly states the assumption associated with this: "Again, this does not assume that the prior distribution is correct, only that the identification of locally dominant sources is correct".

12.) Page 4509, section 5. It may be useful to discuss the different time frame these studies focused their observation-inversions on. The different studies have been conducted focused on different years, and one could speculate that could contribute to the differences (though I find it more likely different observing network/transport/inverse strategy has a bigger role). Miller et al., 2012 and Wecht et al., 2014 had different years of focus.

Excellent question! We have ongoing work investigating the difference between these studies and is the focus of a manuscript that is in preparation. So it's not something we wanted to belabor in this manuscript.

13.) Page 4509 17-18. This is a bit misleading as stated. The Miller study had little to no sensitivity to many regions with stronger wetland sources (such as Florida), and so pre-subtracting wetland emissions is not really an important component of the analysis in that respect (wetland emissions from Florida could be increased by multiple Tg and it would not affect the Miller inversion this pre-subtraction matters, but it is not a simple Tg subtraction from the net).

This seems to be a circular argument. I fail to see how one can claim to have a US methane budget while simultaneously being insensitive to regions that could emit multiple Tg of methane. The Miller et al. paper report an estimate of US methane emissions. Additionally, we confirmed that the wetland source from their work was 2.7 Tg a-1 (Scot M. Miller, personal communication, 31 March 2015).

Further, the wetland regions are not necessarily confined to regions that Miller et al. was insensitive to. Examination of US wetland emissions from the WETCHIMP (Melton et al., 2013; Wania et al., 2013) shows many regions that the Miller et al. study was sensitive to with non-negligible emissions in the WETCHIMP models.

14.) Page 4510, line 15-16. This is a really important statement that needs to be made clear in the abstract and conclusion as well.

We agree that it is an important statement and that's why we listed ranges for all of the attributions in the abstract and throughout the manuscript. Additionally, the final paragraph of the conclusions reiterates this point:

Lines 372-373: "We find that either oil/gas or livestock emissions dominate the correction to prior emissions depending on the assumptions regarding prior errors."

15.) Page 4511 line 24. Suggest change 'We attribute' to something along the lines of: "The model framework attributes (with potentially larger uncertainty)"

See response to major comment 1.

16.) Table 2: What does the increase in all the mean biased post-inversion mean?

See response to comment #12 from Reviewer #2.

17.) Figure 7: There appears to be a plotting problem with the error bars.

The error bars were plotted correctly. As was mentioned in the figure caption, the error bar on the gray bar for the four sectors (oil/gas, livestock, waste, and coal) is from the sensitivity inversion. We have to clarified the figure caption:

Caption: "Error bars on sectoral emissions for our results are defined by the sensitivity inversion with 30% prior uncertainty for all state vector elements."

Reviewer #2 Comments:

This is an interesting and well-written paper. I have what I think are only small concerns that could use some clarification.

Minor Comments:

1.) P4500, L27 – Looking at Figure 1, I don't understand how the relevant spatial differences can be 10 ppb. It looks like they could be 4 times this number.

Fig. 1 shows the “raw” observations. The spatial patterns are, largely, a convolution of the sources and topography. Higher elevation would mean that a larger fraction of the total column is in the stratosphere (where methane concentrations drop off rapidly). Looking at regions of comparable elevation, we see relevant spatial differences of ~10 ppbv.

2.) P4501, L28-29 – This statement doesn't seem to be backed up with evidence. How do we know what the day to day variability is, and how would it follow that GOSAT can constrain the multi-year average? Please elaborate.

Central limit theorem allows us to “beat down the error” with more observations. So the observations may not be accurate enough to constrain a given grid cell but with many observations we can constrain a mean value. We found the seasonal cycle to be reasonably well represented in the prior, so we can use multiple years of observations to constrain the mean. We have updated the text accordingly:

Lines 95-96: “With a mean single-scene instrument precision of 13.3 ppbv, reducible by temporal or spatial averaging, GOSAT...”

3.) P4501, L14-24 – Do we have to worry that Xco2 produced by LMDZ is biased in places other than near large urban sources? The models can produce significant biases, especially in sparsely observed regions. What sort of errors can we expect to arise from this? Could comparisons with full physics retrievals be shown for other latitudes and locations?

This was something that we discussed with other co-authors prior to submission. We tested the impact of tested the impact of potential XCO₂ gradients or biases by simply replacing the model XCO₂ with a full physics XCO₂. So usually we obtain the proxy retrievals as:

$$XCH_4^a = \frac{XCH_4^*}{XCO_2^*} XCO_2^{\text{model}}$$

where the asterisk is a non-scattering retrieval and the XCO₂ comes from a model. However, we can instead get a proxy retrieval using an XCO₂ from a full-physics retrieval:

$$XCH_4^b = \frac{XCH_4^*}{XCO_2^*} XCO_2^{\text{full-physics}}$$

Taking the difference and rearranging gives us:

$$XCH_4^a - XCH_4^b = XCH_4^a \left(1 - \frac{XCO_2^{\text{full-physics}}}{XCO_2^{\text{model}}} \right)$$

We have plotted this difference in the figure below. We can see there are some differences

but we don't see any major systematic biases.

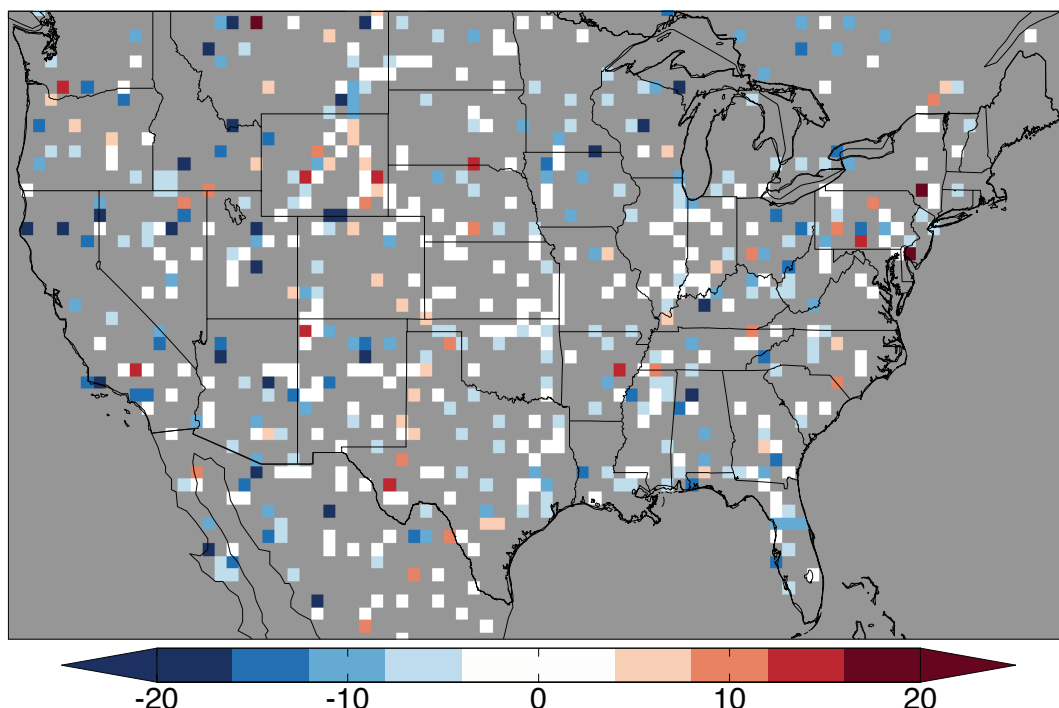


Figure 2: Difference between the proxy retrievals using a modeled CO_2 and the full-physics retrieval for CO_2 : $(\text{XCH}_4^a - \text{XCH}_4^b)$.

We have updated the text to reflect this.

Lines 115-117: “We determined the extent of the bias by replacing XCO_2 in Eq. (1) with (sparser) XCO_2 data from a full-physics GOSAT retrieval. This indicates a 14 ppbv low bias in Los Angeles but much weaker biases on regional scales.”

4.) P4501, L25-27- It seems odd to say that the GOSAT will be compared to surface and aircraft obs. by using a model (and not even one that has emissions optimized with the obs.!). I guess the idea is that GEOS-Chem agrees with the obs, so therefore GEOS-Chem is a proxy for the obs. On the other hand, Figure A3 seems to imply a bias in GEOS-Chem that can be as large 20-25 ppb vs flasks. Is this considered unimportant? If so, please explain because this difference could be as large as the 10 ppb spatial difference mentioned above.

Good question! There is a lot of variability in the comparison of the model with surface flasks. A 10 ppb difference is quite small. We have added 1- σ error bars to the middle panel of A3. The uncertainties of the running medians all bound zero.

5.) Figure A3 - While I'm on the topic of Figure A3, I'd like to suggest that the N-S gradient figures be shown using a vertical axis that ranges from 1700 to 1900. It's hard to see the differences on this scale. The same goes for the middle figure, which could have a vertical axis ranging from -10 to 30 or something like that.

We have compressed the vertical axis of the N-S gradient figures.

6.) P4502, L10 - “Comparisons over North America with NOAA. . .” It’s not clear to me what is being compared here - I think it’s the model, but it could be GOSAT. Please clarify.

We have rephrased this:

Lines 128-129: “Comparing GEOS-Chem at 4°×5° over North America with ...”

7.) P4502, L16-29 - I don’t think I follow this discussion that determines the model stratosphere is the cause of the latitudinal bias between the model and GOSAT. The solar zenith angle is of course determined by latitude (and season), so how can latitude and sza be separated? Also, if it is the model, then how was the stratosphere determined to be the culprit?

Also, why does the model-GOSAT residual still show differences over Greenland (while it doesn’t show differences over the Tibetan Plateau)?

I think it would be very helpful to compare the modeled stratospheric CH₄ to other data, like stratospheric aircraft or air cores.

We touch on this in our response to Minor Comment #8 from Reviewer #1. We tried examining a wealth of observations and have not been able to conclusively attribute the bias to either GOSAT or GEOS-Chem. Solar Zenith Angle and latitude can be separated in a simple linear model. We construct simple linear models for the bias (read: multiple linear regression), used a model selection metric, the Bayesian Information Criterion, and found that linear models with latitude were a better predictor for the bias than SZA. This leads us to suspect the bias is in GEOS-Chem because, physically, SZA is the parameter that would affect the satellite retrieval, not latitude. Again, this is not conclusive and we have been working with Kat M. Saad (Caltech; Wennberg group) using their partial TCCON columns to evaluate the model and interrogate a possible stratospheric bias. However we have not yet found the source of the bias and think it is best to leave this for future investigation.

As for the Greenland question, there are very few observations over Greenland and they are quite noisy (see Fig. 2b).

We have rephrased the bias correction paragraph:

Lines 145-147: “This suggests a potential bias in the GEOS-Chem simulation of methane in the polar stratosphere, which warrants further investigation with observations such as TCCON partial columns (Saad et al. 2014; Wang et al. 2014). In any case, we remove the bias using...”

8.) Sections 3, 4 - It would be helpful to be more explicit about what is being estimated. Are parameters in the global inversion grid-box total emissions? Are they time dependent? Or annual totals?

We define the state vector (what is being estimated) in the first paragraph of the global and North American inversion sections:

Lines 159-160: “The state vector consists of scaling factors for emissions at $4^{\circ} \times 5^{\circ}$ resolution for June 2009–December 2011.”

Lines 208-210: “We only solve for the spatial distribution of emissions, assuming that the prior temporal distribution is correct (aseasonal except for wetlands and open fires, see Appendix).”

9.) P4504, L28-29- I don’t understand this statement. How does one trust the prior distributions of sources processes at a local scale, but not at global scales?

It’s not that we trust the local distributions of sources more than the global scales but we trust the relative numbers more. This is far from ideal but we feel it’s the best we can do given the current observing system. Ideally we would use observations of correlated species (e.g., ethane or $\delta^{13}\text{C}$) or a better prior. We elaborate on this in the conclusions:

Lines 373-375: “This limitation [source attribution] could be addressed in the future through better specification of the prior source distribution using high-resolution information on activity rates, and through the use of correlated variables in the inversion.”

10.) P4505, L21-26 - I like this approach of reducing parameters, but these lines could use a bit more elaboration. What is the rationale for using the adjustments from the global inversions to cluster grid boxes?

Ah, that was a single iteration from an adjoint-based inversion at high resolution. The reason for including that parameter was to account for sources that may be present in the satellite data and missing from other “criteria”. We elaborate on this in the companion paper (Turner and Jacob, 2015). We have also updated the text accordingly:

Lines 212-221: “The RBFs are constructed from estimation of the factors driving error correlations between the native resolution state vector elements including spatial proximity, correction from one iteration of an adjoint-based inversion at $\frac{1}{2}^{\circ} \times \frac{2}{3}^{\circ}$ resolution, and prior source type distributions. Including the correction from the adjoint-based inversion allows us to account for sources not included in the prior.”

11.) P4508, L16-17 - Here again is the statement about prior source distributions that I don’t understand. How can we trust the allocation to different sources, but not the spatial distribution?

See response to comment #9.

12.) Table 2 - Why is the posterior mean bias larger for the posterior than the prior? Shouldn’t the fit improve after inversion?

These observations are independent from the inverse modeling framework. Thus, there won't necessarily be an improvement. The posterior did improve the comparison to the independent observations by most metrics we examined.

Estimating global and North American methane emissions with high spatial resolution using GOSAT satellite data

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Abstract. We use 2009–2011 space-borne methane observations from the Greenhouse Gases Observing SATellite (GOSAT) to ~~constrain-estimate~~ global and North American ~~inversions-of~~ methane emissions with $4^\circ \times 5^\circ$ and up to $50\text{ km} \times 50\text{ km}$ spatial resolution, respectively. ~~The-GEOS-Chem~~ and GOSAT data are first evaluated with atmospheric methane observations from surface ~~and tower~~ networks (NOAA/~~ESRL~~, TCCON) and aircraft (NOAA/~~DOE~~ESRL, HIPPO), using the GEOS-Chem chemical transport model as a platform to facilitate comparison of GOSAT with in situ data. This identifies a high-latitude bias between the GOSAT data and GEOS-Chem that we correct via quadratic regression. ~~The-surface-and-aircraft-data-are-subsequently-used-for-independent-evaluation-of-the-methane-source-inversions.~~ Our global adjoint-based inversion yields a total methane source of 539 Tg a^{-1} ~~and points to a large East Asian overestimate in~~ with some important regional corrections to the EDGARv4.2 inventory used as a prior. Results serve as dynamic boundary conditions for an analytical inversion of North American methane emissions using radial basis functions to achieve

high resolution of large sources and provide ~~full~~ error characterization. We infer a US anthropogenic methane source of $40.2\text{--}42.7\text{ Tg a}^{-1}$, as compared to $24.9\text{--}27.0\text{ Tg a}^{-1}$ in the EDGAR and EPA bottom-up inventories, and $30.0\text{--}44.5\text{ Tg a}^{-1}$ in recent inverse studies. Our estimate is supported by independent surface and aircraft data and by previous inverse studies for California. We find that the emissions are highest in the South-Central US, the Central Valley of California, and Florida wetlands, large isolated point sources such as the US Four Corners also contribute. ~~We~~ Using prior information on source locations we attribute 29–44 % of US anthropogenic methane emissions to livestock, 22–31 % to oil/gas, 20 % to landfills/~~waste-water~~ wastewater, and 11–15 % to coal ~~with~~. Wetlands contribute an additional $9.0\text{--}10.1\text{ Tg a}^{-1}$ ~~source from wetlands~~.

1 Introduction

Methane (CH_4) ~~is the most important anthropogenic greenhouse gas after carbon dioxide~~ emissions have contributed 0.97 W m^{-2} in global radiative forcing of climate since pre-industrial times, second only to CO_2 with 1.7 W m^{-2} (IPCC, 2013). As a short-lived climate forcing agent (lifetime ~ 10 years), ~~it~~ methane may provide a lever for slowing near-term climate change (Ramanathan and Xu, 2010; Shindell et al., 2012). Major anthropogenic sources include natural gas and petroleum production and use, coal mining, waste (landfills and wastewater treatment), livestock, and rice cultivation. Wetlands are the largest natural source. The present-day global emission of methane is $550 \pm 60\text{ Tg a}^{-1}$, constrained by knowledge of the global tropospheric hydroxyl radical (OH) concentration from the methylchloroform budget (Prather et al., 2012). However, allocation by source types and regions is very uncertain (Dlugokencky et al., 2011). Here we use GOSAT space-borne observations for 2009–2011 to improve our understanding of global and North American methane emissions using a high-resolution inversion technique ~~(?)~~ (Turner and Jacob, 2015).

The US Environmental Protection Agency (EPA) produces national emission inventories for anthropogenic methane, with a total of 27.0 Tg a^{-1} in 2012 including 34 % from livestock, 29 % from oil/gas extraction and use, 21 % from waste ~~(landfills and waste-water)~~, and 11 % from coal mining (EPA, 2014). Inverse studies using observations of atmospheric methane concentrations suggest that the EPA inventory may be too low by up to a factor of 2, although they differ as to the magnitude and cause of the underestimate (Katzenstein et al., 2003; Kort et al., 2008; Xiao et al., 2008; Karion et al., 2013; Miller et al., 2013; Wecht et al., 2014a; Caulton et al., 2014). There is strong national and international interest in regulating methane emissions (President’s Climate Action Plan, 2013, 2014; Climate and Clean Air Coalition, 2014), particularly in the context of increasing natural gas exploitation and fracking, but uncertainty in the emission inventory makes regulation problematic.

Space-borne observations of atmospheric methane concentrations in the short wave infrared (SWIR) are a unique resource for constraining methane emissions because of the dense and continuous data that they provide. SWIR instruments measure column concentrations with near-uniform ver-

tical sensitivity down to the surface. Data are available from the SCIAMACHY instrument for 2003–2012 (Frankenberg et al., 2005, 2011) and from the TANSO-FTS instrument aboard GOSAT for 2009–present (Butz et al., 2011; Parker et al., 2011; hereafter we refer to the instrument as “GOSAT”). GOSAT has higher precision and pixel resolution than SCIAMACHY (0.6 % and 10 km × 10 km vs. 1.5 % and 30 km × 60 km) but the observations are not as dense. The GOSAT retrievals are in good agreement with surface-based column measurements (Parker et al., 2011; Butz et al., 2011; Schepers et al., 2012; Fraser et al., 2013; Monteil et al., 2013; Cressot et al., 2014; Alexe et al., 2015).

Previous inversions of methane emissions using satellite data have mainly focused on the global scale, optimizing emissions with coarse spatial resolution (Bergamaschi et al., 2007, 2009, 2013; Fraser et al., 2013; Monteil et al., 2013; Cressot et al., 2014; Alexe et al., 2015). This limits the interpretation of the results because emissions from different source types have large spatial overlap (Fung et al., 1991). [Spatial overlap is reduced at higher resolution, thus optimizing emissions at high spatial resolution can help improve source attribution.](#) Wecht et al. (2014a) used SCIAMACHY data for July–August 2004 in a higher-resolution ($\sim 100 \text{ km} \times 100 \text{ km}$) inversion of methane emissions in North America, but they were unable to achieve such a resolution using GOSAT because of the sparser data (Wecht et al., 2014b).

Here we use three years (2009–2011) of GOSAT data to constrain global and North American methane emissions with high spatial resolution, exploiting both the longer record and a new analytical inversion method where the state vector of emissions is defined optimally from a set of radial basis functions (?) (Turner and Jacob, 2015). We begin by evaluating the GOSAT retrievals with surface, aircraft, and total column observations using the GEOS-Chem chemical transport model (CTM; described in Appendix) as an intercomparison platform. This identifies a high-latitude bias between GOSAT and GEOS-Chem that we correct. We then use GOSAT observations to constrain global methane sources with the GEOS-Chem model and its adjoint at $4^\circ \times 5^\circ$ resolution, and apply the results as boundary conditions to optimize North American methane sources with up to $50 \text{ km} \times 50 \text{ km}$ resolution and full-error characterization.

2 GOSAT observations and bias correction

GOSAT (Kuze et al., 2009) was launched in January 2009 by the Japan Aerospace Exploration Agency (JAXA). Methane abundance is determined by analysis of the spectrum of backscattered solar radiances in the SWIR near $1.6 \mu\text{m}$. Data are available from April 2009 on. GOSAT is in Sun-synchronous low earth orbit with an equator overpass of 12:45–13:15 LT. The instrument observes five cross-track nadir pixels (three cross-track pixels after August 2010) with a footprint diameter of 10.5 km, a cross-track spacing of about 100 km, an along-track spacing of 90–280 km, and a 3 day revisit time. We use the version 4 proxy methane retrievals from Parker et al. (2011) that

pass all quality flags (<http://www.leos.le.ac.uk/GHG/data/styled/index.html>). The retrievals provide a weighted column average dry-mole fraction of CH_4 , X_{CH_4} , with averaging kernels to describe the vertical weighting. The averaging kernels show near-uniform vertical sensitivity in the troposphere and decreasing sensitivity above the tropopause (see Butz et al., 2010). The estimated single-retrieval precision is scene-dependent and averages 13.3 ppbv or 0.8 % (Parker et al., 2011).

Figure 1 shows the mean methane concentrations for June 2009–December 2011 observed by GOSAT and used in this work. There are 590 675 global observations including 74 687 for the North American window of our high-resolution inversion. The GOSAT sampling strategy of consistently revisiting the same locations provides a high density of observations over the sampled locations but the coverage is not continuous (gray areas in Fig. 1). There are data over oceans from Sun glint retrievals (Butz et al., 2011) but not in the Parker et al. (2011) product [used here](#). Methane concentrations are highest over East Asia where rice, livestock, and fossil fuels contribute large sources. They are also high over the eastern US. Low concentrations over elevated terrain (Tibetan Plateau, western US) reflect in part a larger relative contribution of the stratosphere to the column-average mixing ratio. We see from Fig. 1 that relevant spatial differences in methane mixing ratio for our inversion are of the order of 10 ppbv. With a mean single-scene instrument precision of 13.3 ppbv, [reducible by temporal or spatial averaging](#), GOSAT cannot resolve day-to-day variability of emissions but can strongly constrain a multi-year average.

Previous studies have validated the GOSAT data with surface-based FTIR methane column abundances from the Total Carbon Column Observing Network (TCCON; Wunch et al., 2011). These studies have generally found GOSAT retrievals to be within their stated precisions (Parker et al., 2011). Schepers et al. (2012) pointed out that a full validation of the GOSAT retrievals would require a more extensive validation network than is available from TCCON. Satellite observations by solar backscatter tend to be subject to high-latitude biases because of large solar zenith angles [resulting in longer path lengths and higher interference with clouds](#). Monteil et al. (2013) did not include a latitudinal bias correction in their inverse analysis of GOSAT data, but Cressot et al. (2014) used a bias correction based on the geometric air mass factor and Fraser et al. (2013) added a latitudinal bias correction that was fitted as part of the inversion.

The Parker et al. (2011) retrieval uses CO_2 as a proxy for the lighthpath to minimize common spectral artifacts from aerosol scattering and instrument effects (Frankenberg et al., 2005; Butz et al., 2010):

$$X_{\text{CH}_4} = \frac{X_{\text{CH}_4}^*}{X_{\text{CO}_2}^*} X_{\text{CO}_2} \quad (1)$$

where $X_{\text{CH}_4}^*$ and $X_{\text{CO}_2}^*$ are the dry-air mole fractions retrieved from GOSAT under the assumption of a non-scattering atmosphere and X_{CO_2} is the column-average dry-air mole fraction of CO_2 , estimated from the LMDZ global CTM with $3.75^\circ \times 2.5^\circ$ spatial resolution. This could lead to localized bias in areas of concentrated CO_2 sources. We determined the extent of the bias by replacing X_{CO_2}

in Eq. (1) with (sparser) X_{CO_2} data from a full-physics GOSAT retrieval. This indicates a 14 ppbv
120 low bias in Los Angeles but much weaker biases on regional scales.

Here we examined the consistency of GOSAT with a large body of independent surface and air-
craft measurements of methane concentrations by using the GEOS-Chem CTM with prior methane
emissions (Table 1 and Figs. A1 and A2) as an intercomparison platform. Table 2 gives summary
comparison statistics and more details are in the Appendix (Figs. A3–A5). Global comparisons with

125 HIPPO II-V aircraft profiles across the Pacific (<http://hippo.ornl.gov>; Wofsy, 2011), the NOAA co-
operative flask sampling network (<http://www.esrl.noaa.gov/gmd/ccgg/flask.php>), and the TCCON
network (<http://tecon.ornl.gov>, GGG2012 version; Washenfelder et al., 2006; Deutscher et al., 2010; Wunch et al., 2010, 2011; Mes

that GEOS-Chem accurately simulates the global features of the methane distribution including the
meridional gradient in different seasons with no significant bias across multiple years and seasons

130 (Figs. A3–A5). One would then expect similar agreement in the comparison of GEOS-Chem to

GOSAT. ~~Comparisons~~ Comparing GEOS-Chem at $4^\circ \times 5^\circ$ over North America with the NOAA
flasks TCCON, the NOAA/ESRL Global Greenhouse Gas Reference Network (surface flasks, tall
tower network, TCCON, and vertical profiles from the NOAA/DOE aircraft program (<http://www.esrl.noaa.gov/gmd/ccgg/aircraft/>; E

135 weaker correlations (~~$R^2 = 0.56$~~ $R^2 = 0.40$ – 0.57 , 0.60) and the reduced-major-axis regression slopes
(~~0.64 – 0.67~~ 0.71 – 0.75) suggest a $\sim 30\%$ prior underestimate of North American emissions. Reduction
of this bias will provide an independent check on our inversion results.

Figure 2 compares the GOSAT methane observations (X_{CH_4}) to GEOS-Chem values sampled at
the location and time of the observations, and with local averaging kernels applied. There is a lati-
tudinal background pattern in the difference between GEOS-Chem and GOSAT. The bias becomes
140 significant at latitudes poleward of 50° . Since GEOS-Chem is unbiased in its simulation of the tro-
pospheric meridional gradient relative to the surface and aircraft data (Table 2, Fig. A3), we attribute
the high-latitude bias to errors in either the: GOSAT retrieval or GEOS-Chem stratospheric methane.

Bias corrections that are a function of latitude or air mass factor (solar zenith angle) should be able

145 to correct for this. However, a bias in the GOSAT data would be expected to correlate better with the
air mass factor while a bias in the model stratosphere may correlate better with latitude. We find lat-
itude to be a better bias predictor based on the Bayesian information criterion (quadratic regression
in Fig. 2c). This suggests a potential bias in the GEOS-Chem simulation of methane in the polar
stratosphere. ~~We,~~ which warrants further investigation with observations such as TCCON partial

150 columns (Saad et al., 2014; Wang et al., 2014). In any case, we remove the bias using the quadratic
regression and Fig. 2d shows the resulting mean differences between GEOS-Chem and GOSAT after
this bias correction. The differences point to errors in the GEOS-Chem prior emissions that we will
correct in the inversion.

3 Global inversion of methane emissions

We use the bias-corrected GOSAT data to infer global methane emissions at $4^\circ \times 5^\circ$ resolution ~~by using-with~~ an adjoint-based four-dimensional variational data assimilation system (Henze et al., 2007; Wecht et al., 2012, 2014a) ~~to minimize~~. The system minimizes a cost function (\mathcal{J}) with Gaussian errors,

$$\mathcal{J}(x) = \frac{1}{2}(\mathbf{y} - \mathbf{K}x)^T \mathbf{S}_O^{-1}(\mathbf{y} - \mathbf{K}x) + \frac{1}{2}(x_a - x)^T \mathbf{S}_a^{-1}(x_a - x) \quad (2)$$

Here x_a is the vector of prior emissions (see Table 1 and Fig. A1), \mathbf{y} is the vector of GOSAT observations, $\mathbf{K} = \partial \mathbf{y} / \partial x$ is the Jacobian matrix of the GEOS-Chem methane simulation used as forward model, and \mathbf{S}_a and \mathbf{S}_O are the prior and observational error covariance matrices respectively.

The state vector consists of scaling factors for emissions at $4^\circ \times 5^\circ$ resolution for June 2009–December 2011. The prior emissions are mainly from the EDGARv4.2 inventory for anthropogenic sources (European Commission, 2011), and Pickett-Heaps et al. (2011) for wetlands. Table 1 gives a summary and further details are in the Appendix. The error covariance matrices are taken to be diagonal, implying no error correlation on the $4^\circ \times 5^\circ$ grid. We assume 50 % error variance on the prior for $4^\circ \times 5^\circ$ grid cells as in Monteil et al. (2013).

Observational error variances are estimated following Heald et al. (2004) by using residual ~~SDs~~ standard deviations of the differences between observations and the GEOS-Chem simulation with prior emissions, as shown in Fig. 2b. This method attributes the mean bias on the $4^\circ \times 5^\circ$ grid to errors in emissions (to be corrected by the inversion) and the residual error to observational error (including contributions from instrument retrieval, representation, and model transport errors). If the resulting observational error variances are less than the local retrieval error variances reported by Parker et al. (2011) then the latter are used instead. This is the case for 58 % of the observations, implying that the observational error is dominated by the instrument retrieval error.

The GEOS-Chem forward model and its adjoint are as described by Wecht et al. (2014a). We optimize methane emissions from 1 June 2009 to 1 January 2012. The forward model is initialized on 1 January 2009 with concentrations from Wecht et al. (2014a). There is no significant global bias in the simulation, as discussed above. The 5-month spin-up allows for the establishment of gradients driven by synoptic motions and effectively removes the influence of the initial conditions.

Figure 3 shows the prior and posterior 2009–2011 emissions. We evaluated the posterior emissions in a GEOS-Chem forward simulation by comparison with the global independent observational datasets of Table 2. The prior simulation showed high correlation and little bias. The posterior simulation shows similar results. The increase in mean bias relative to the TCCON data is not significant. As pointed out above, the global data sets mainly test the global emissions and large-scale meridional gradients. ~~The inversion~~ Since we used them previously to justify a bias correction in the comparison between GEOS-Chem and GOSAT, they do not provide a true independent test of the

inversion results. Nevertheless, we see that the inversion does not degrade the successful simulation

of the background meridional gradient in the prior GEOS-Chem simulation.

The total posterior methane emission is 539 Tg a^{-1} , unchanged from the prior (537 Tg a^{-1}). This source is within the $548^{+21}_{-22} \text{ Tg a}^{-1}$ range of current estimates reported by Kirschke et al. (2013) and IPCC (2013). However, we find large regional differences compared to the prior. Emissions from China are revised downward by 50 % from 29.2 to 14.7 Tg a^{-1} , consistent with Bergamaschi et al. (2013) who find that EDGARv4.2 Chinese coal emissions are too large. This overestimate in Chinese methane emissions is also seen by Bruhwiler et al. (2014) who assimilated the 2000–2010 NOAA surface observations into CarbonTracker using an ensemble Kalman filter. Emissions in India are also too high, while emissions in Southeast Asia are ~~10 too low due to underestimated rice emissions too low~~. Emissions from wetlands in central Africa are too high. Emissions in northern South America are too low. Corrections in North America are discussed in the next section.

We inferred the contributions from different source types to our posterior emissions by assuming that the prior inventory correctly partitions the methane by source type (see Appendix and Table 1) in each $4^\circ \times 5^\circ$ grid cell. This does not assume that the global distribution of source types is correct in the prior, only that the local identification of dominant sources is. We find only modest changes to the global partitioning by source types, with the exception of coal and rice, partly reflecting regional offsets. For example, ~~the global~~ wetland emissions increase globally by only 5 Tg a^{-1} but decrease by 24 Tg a^{-1} in the African wetlands ; ~~this regional decrease is partly offset by a~~ while increasing by 10 Tg a^{-1} ~~increase~~ in northern South America.

4 North American inversion of methane emissions

We optimize methane emissions over North America by using the nested GEOS-Chem simulation at $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ horizontal resolution ($\sim 50 \text{ km} \times 50 \text{ km}$) over the North American window in Fig. 1. Time-dependent boundary conditions for this nested simulation are from the global model at $4^\circ \times 5^\circ$ horizontal resolution using the posterior emissions derived above. We only solve for the spatial distribution of emissions, assuming that the prior temporal distribution is correct (aseasonal except for wetlands and open fires, see Appendix).

Following Turner and Jacob (2015), the dimension of the emissions state vector for the nested North American inversion is optimally reduced from the native $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution ($n = 7366$) in order to (1) improve the observational constraints on individual state vector elements and (2) enable an analytical inversion with full error characterization. This is done by aggregating similar state vector elements with a Gaussian mixture model (Bishop, 2007). We find that an optimal reduction with negligibly small aggregation error can be achieved using 369 radial basis functions (RBFs) with Gaussian kernels. The RBFs are constructed from qualitative knowledge estimation of the factors driving ~~prior error correlations in~~ error correlations between the native resolution state

vector elements including spatial proximity, ~~initial correction from~~ correction from one iteration of
 225 an adjoint-based inversion at $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution, and prior source type distributions. Including the
correction from the adjoint-based inversion allows us to account for sources not included in the
prior. Each $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ native-resolution grid square is projected onto an aggregated state vector using
 the RBFs. This preserves native resolution where needed (in particular for large point sources) and
 aggregates large regions where emissions are uniformly small.

230 Our optimal estimate of North American emissions was obtained by analytical solution to Eq. (2)
 (cf. Rodgers, 2000), using the Jacobian matrix \mathbf{K} constructed column by column for the aggregated
 state vector. This analytical approach provides the posterior covariance matrix $\hat{\mathbf{S}}$ and averaging ker-
 nel matrix \mathbf{A} as part of the solution and thus fully characterizes the errors and information content
 of the inversion results.

235 The observational error covariance matrix is assumed diagonal with terms specified as the larger
 of the residual error variance and the retrieval error variance, same as for the global inversion. The
 prior error covariance matrix is assumed diagonal because the radial basis functions are designed
 to capture spatial correlations in the emissions. We assume 100 % error on emissions at the native
 $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution. For RBFs encompassing larger spatial regions we assume that the error is reduced
 240 following the central limit theorem:

$$S_{a,\{i,i\}} = \frac{s_a}{\sqrt{\sum_j w_{i,j}}} \quad (3)$$

where $S_{a,\{i,i\}}$ is the ~~i th~~ ^{i th} diagonal of \mathbf{S}_a , s_a is the prior uncertainty at the native resolution (100 %),
 and the summation is for the weights of the i th RBF over all $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ grid squares (index j). This error
 reduction assumes that the errors on the native-resolution grid cells are independent and identically
 245 distributed, which may be overly optimistic. We examined the sensitivity to this assumption by
 conducting an alternate inversion with a relative error of 30 % for all state vector elements, similar to
 the approach taken by Wecht et al. (2014a) using a hierarchical clustering method for the state vector.

Figure 4 shows the prior and posterior 2009–2011 emissions. Total posterior emissions in North
 America (Table 1) are 44 % higher than the prior, with large increases in the South-Central US and
 250 weak decreases for the Canadian wetlands. Contiguous US emissions are 52 Tg a^{-1} , 70 % higher
 than the prior. The broad correction patterns are consistent with the coarse global results in Fig. 3
 that used a completely different inversion method. Our sensitivity inversion assuming 30 % prior
 error on all state vector elements yields the same North American and contiguous US totals to within
 3 %.

255 We evaluated the posterior emissions in a GEOS-Chem simulation over North America by com-
 parison to the independent observations from Table 2. We find great improvement in the ability of the
 model to reproduce these observations, as illustrated by the scatterplots of Fig. 5. The reduced-major-
 axis (RMA) regression slopes improve from ~~0.71 to 0.92~~ 0.72 to 1.03 for the NOAA/~~DOE aircraft~~

~~profiles and from 0.64 to 1.00~~ ESRL tall tower network, from 0.75 to 0.94 for the NOAA/ESRL aircraft profiles, and from 0.67 to 1.01 for the NOAA ~~cooperative surface flask network~~ surface flasks.

Another independent evaluation of our posterior emissions is the estimate for California. California's methane emissions have been extensively studied with aircraft and ground-based observations over the past few years in order to address statewide greenhouse gas regulation targets (Zhao et al., 2009; Wunch et al., 2009; Hsu et al., 2010; Peischl et al., 2012; Wennberg et al., 2012; Jeong et al., 2012, 2013; Peischl et al., 2013; Santoni et al., 2014; Wecht et al., 2014b). Figure 6 shows that our posterior emissions are 20 % higher than the EDGARv4.2 prior inventory for the state of California and 50 % lower for the Southern California Air Basin (SoCAB). Other studies constrained with dense aircraft and ground-based observations are consistent with ours. Our estimate for SoCAB could be biased low due to an underestimate of local CO₂ in the GOSAT retrieval (see Sect. 2). Wecht et al. (2014b) previously found that GOSAT observations were not sufficiently dense to constrain methane emissions in California. However, they only used a 2 month record and tried to constrain emissions at $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution, incurring large smoothing error. By using a longer time record and an optimally defined state vector we ~~find that GOSAT can effectively constrain emissions at the state scale~~ achieve much better success.

Figure 4 (top right panel) shows the averaging kernel sensitivities for the North American methane emission inversion, defined as the diagonals of the averaging kernel matrix. The inversion has 39 degrees of freedom for signal (DOFs), meaning that we can exactly constrain 39 pieces of information in the distribution of methane emissions. ~~In fact, this~~ This information is spread over the continent and mixed with prior constraints as described by the averaging kernel matrix. We can use the averaging kernel sensitivities in Fig. 4 to determine which regions are most responsive to the inversion. These include California, the Canadian wetlands, and the Southeast and Central US. Large isolated point sources such as the US Four Corners (a large source of coalbed methane at the corner of Arizona, New Mexico, Colorado, and Utah) are also strongly sensitive to the inversion.

We see from Fig. 4 that the prior underestimate of North American methane emissions is largely due to the Central US, the Canadian Oil Sands, Central Mexico, California, and Florida. Various large point sources such as the US Four Corners also contribute. We also find regions where the prior is too high including the Hudson Bay Lowlands, SoCAB, and parts of Appalachia. This suggests that oil/gas and livestock emissions are higher than given in EDGARv4.2 while coal emissions are lower. The overestimate in SoCAB is likely because EDGARv4.2 uses urban and rural population as a spatial proxy for landfills and waste water (Wunch et al., 2009). The underestimate in Florida is most likely due to wetland sources.

As with the global inversion, we infer the contributions from different methane source types by assuming that the prior inventory correctly attributes the source types in a given $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ grid cell. Again, this does not assume that the prior distribution is correct, only that the identification of locally dominant sources is correct. Results are shown in Fig. 7. We see that the increase relative to the prior

is mainly driven by anthropogenic sources. This can be compared to the US EPA anthropogenic inventory (EPA, 2014), which is based on more detailed bottom-up information than EDGARv4.2 but is only available as a national total. We find an anthropogenic source for the contiguous US of 40.2–42.7 Tg a⁻¹, as compared to 27.0 Tg a⁻¹ in the US EPA inventory. The largest differences are for the oil/gas and livestock sectors. Depending on the assumptions made regarding the prior error, oil/gas emissions from our inversion are 13–74 % higher than the EPA estimate and contribute 17–26 % of contiguous US methane emissions. Livestock emissions are 36–85 % higher than the EPA estimate and contribute 24–33 % of contiguous US methane emissions. Waste and coal emissions are also higher in our posterior estimate than in the EPA inventory.

5 Comparison to previous inverse studies

Several past inverse analyses have estimated methane emissions in the contiguous US with differing conclusions, in particular the work of Miller et al. (2013) and Wecht et al. (2014a). Miller et al. (2013) used in situ observations for 2007–2008 from ground stations and aircraft. They found the EPA inventory to be underestimated by a factor of 1.5 nationally with the largest underestimates in fossil fuel source regions. This is in contrast to Wecht et al. (2014a) who used July–August 2004 observations from SCIAMACHY. They found that the EPA inventory was underestimated by only 10 %, with the major discrepancy being livestock emissions underestimated by 40 %.

Our continental-scale inversion yields a total US methane emissions of 52.4 Tg a⁻¹ and anthropogenic source of 42.8 Tg a⁻¹. The general spatial pattern of the posterior emissions is similar to those of Miller et al. (2013) and Wecht et al. (2014a) but the total methane emissions found here are more similar to Miller et al. (2013) who ~~find~~found US total and anthropogenic emissions of 47.2 and 44.5 Tg a⁻¹. The corresponding values obtained by Wecht et al. (2014a) are 38.8 and 30.0 Tg a⁻¹, significantly lower.

Our work finds a larger natural methane source in the contiguous US than Miller et al. (2013), who used a fixed prior wetland source of 2.7 Tg a⁻¹ that was subtracted from the measurements. Our prior and posterior emissions are 5.9 and 9.0–10.1 Tg a⁻¹, respectively, mostly located in Louisiana and Florida and more consistent with Wecht et al. (2014a). Quantifying the wetlands source is important because it subtracts from the anthropogenic source estimate inferred from the inversion. In particular, our anthropogenic source of methane in the contiguous US would be larger than that of Miller et al. (2013) if we had not corrected for the larger wetland source.

Kort et al. (2014) found the Four Corners to be the largest single methane source in the continental US (0.59 Tg a⁻¹) on the basis of SCIAMACHY observations and TCCON observations, with a magnitude 3.5 times larger than EDGARv4.2 and 1.8 times larger than reported by the US EPA Greenhouse Gas Reporting Program (EPA, 2014). This is in contrast to Miller et al. (2013) who found the US Four Corners to be overestimated in EDGARv4.2 but only had weak constraints for

that region. Our work finds methane emissions from the US Four Corners to be $0.45\text{--}1.39\text{ Tg a}^{-1}$ and 3–9 times larger than in the EDGARv4.2 inventory, consistent with the finding of Kort et al. (2014).

Miller et al. (2013) attributed most of the underestimate in the US EPA methane inventory to fossil fuel while Wecht et al. (2014a) attributed it to livestock. We find in our inversion that the source attribution is highly dependent on the specification of the prior error covariance matrix, as shown in Fig. 7. Our standard inversion that adjusts the prior error for the RBF weights (Eq. 3) attributes 31 % of US anthropogenic emissions to oil/gas and 29 % to livestock, so that most of the EPA underestimate is for oil/gas. However, an inversion without this prior error adjustment (error bars in Fig. 7) finds the underestimate to be mainly from livestock. This is because the RBFs associated with livestock emissions tend to cover larger areas of correlated emissions than the point sources associated with oil/gas. An inversion with equal error weighting for different state vector elements will tend to favor correction of the larger elements associated with livestock. With current prior knowledge it is thus difficult to conclusively attribute the US EPA underestimate to oil/gas or livestock emissions. This limitation could be addressed by a better prior knowledge of the spatial distribution of source types or by the use of correlative information (e.g., observations of ethane originating from oil/gas) in the inversion.

6 Conclusions

We used 31 months of GOSAT satellite observations of methane columns (June 2009–December 2011) to constrain methane emissions at high spatial resolution in North America with an inversion based on the GEOS-Chem chemical transport model. We first conducted a global adjoint-based inversion at $4^\circ \times 5^\circ$ resolution and used the resulting optimized fluxes as dynamic boundary conditions for a nested inversion with resolution up to $50\text{ km} \times 50\text{ km}$ over North America.

We began by evaluating the GOSAT observations with a large ensemble of aircraft and surface data (HIPPO, NOAA/ESRL surface flasks, NOAA/DOE-ESRL aircraft, TCCON), using GEOS-Chem as an intercomparison platform. This revealed a high-latitude bias in the ~~GOSAT data~~ GEOS-Chem polar stratospheric methane (or possibly in the ~~GEOS-Chem polar stratospheric methane~~ GOSAT data) that we corrected for the purpose of the inversion. The aircraft and surface data were subsequently used as an independent check of our inversion results.

Our global GOSAT inversion finds a total methane source of 539 Tg a^{-1} with 39 % from wetlands, 22 % from livestock, 12 % from oil/gas, 12 % from waste, 8 % from rice, and 6 % from coal. Comparison to the EDGARv4.2 inventory used as a prior for the inversion indicates that Chinese coal emissions are a factor of 2 too large, consistent with the findings of Bergamaschi et al. (2013) and Bruhwiler et al. (2014). We find large regional corrections to the EDGARv4.2 inventory includ-

ing a 10 Tg a^{-1} increase of the wetland emissions in South America and a 10 Tg a^{-1} increase of rice emissions in Southeast Asia.

Our North American continental-scale inversion used an emission state vector optimally defined with radial basis functions (RBFs) to enable analytical inversion with full error characterization while minimizing aggregation error (Turner and Jacob, 2015). In this manner we could resolve large point sources at a resolution of up to $50 \text{ km} \times 50 \text{ km}$ while aggregating regions with weak emissions. Our posterior anthropogenic methane source for the contiguous US is $40.2\text{--}42.7 \text{ Tg a}^{-1}$, compared to 25.0 Tg a^{-1} in EDGARv4.2 and 27.9 Tg a^{-1} in the US EPA national inventory. Differences are particularly large in the South-Central US. Our posterior inventory is more consistent with independent surface and aircraft data and with previous studies in California. We On the basis of prior emission patterns we attribute 22–31 % of US anthropogenic methane emissions to oil/gas, 29–44 % to livestock, 20 % to waste, and 11–15 % to coal. There is in addition a $9.0\text{--}10.1 \text{ Tg a}^{-1}$ wetlands source.

Our work confirms previous studies pointing to a large underestimate in the US EPA methane inventory. This underestimate is attributable to oil/gas and livestock emissions, but quantitative separation between the two is difficult because of spatial overlap and limitations of the observing system and prior estimates. We find that either oil/gas or livestock emissions dominate the correction to prior emissions depending on assumptions regarding prior errors. This limitation could be addressed in the future through better specification of the prior source distribution using high-resolution information on activity rates, and through the use of correlated variables in the inversion.

Appendix: GEOS-Chem description and evaluation with independent data

We use the v9-01-02 GEOS-Chem methane simulation (<http://acmg.seas.harvard.edu/geos/index.html>; Wecht et al., 2014a) driven by Goddard Earth Observing System (GEOS-5) assimilated meteorological data for 2009–2011 from the NASA Modeling and Assimilation Office (GMAO). The GEOS-5 data have a native horizontal resolution of $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ with 72 pressure levels and 6 h temporal resolution (3 h for surface variables and mixing depths). The results presented here are from nested simulations at the native $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution over North America ($10\text{--}70^\circ \text{ N}$, $40\text{--}140^\circ \text{ W}$) and global simulations at $4^\circ \times 5^\circ$ resolution. Dynamic boundary conditions for the nested simulations are obtained from the global simulations. Methane loss is mainly by reaction with the OH radical. We use a 3-D archive of monthly average OH concentrations from Park et al. (2004). The resulting atmospheric lifetime of methane is 8.9 years, consistent with the observational constraint of 9.1 ± 0.9 years (Prather et al., 2012).

Prior 2009–2011 emissions for the GEOS-Chem methane simulation are from the EDGARv4.2 anthropogenic methane inventory (European Commission, 2011), the wetland model from Kaplan (2002) as implemented by Pickett-Heaps et al. (2011), the GFED3 biomass burning inven-

400 tory (van der Werf et al., 2010), a termite inventory and soil absorption from Fung et al. (1991),
and a biofuel inventory from Yevich and Logan (2003). Wetlands emissions vary with local temper-
ature, inundation, and snow cover. Open fire emissions are specified with 8 h temporal resolution.
Other emissions are assumed aseasonal. Table 1 lists global, North American, and contiguous US
emissions. Figures A1 and A2 show the spatial distributions of the global and North American prior
405 emissions for the five largest source types.

We evaluated GEOS-Chem with surface-based (NOAA/[ESRL](#), TCCON), ~~tower~~ ([NOAA/ESRL](#)),
and aircraft (HIPPO, NOAA/~~DOE~~[ESRL](#)) observations of methane concentrations for 2009–2011,
both as indirect validation of the GOSAT data and as an independent check on our inversion results.
See main text for references for these observations. We convolve GEOS-Chem with the TCCON
410 averaging kernels and priors before comparison with TCCON observations. Figure A3 uses ob-
servations from the NOAA cooperative flask network and from the HIPPO data across the Pacific
to evaluate the global burden and latitudinal gradient in GEOS-Chem. Figure A4 uses observa-
tions from the ~~NOAA cooperative flask network, the NOAA/DOE aircraft program,~~ [ESRL Global
Greenhouse Gas Reference Network](#) and the TCCON column network for a more specific evalua-
415 tion of the model over North America. Figure A5 shows corresponding scatterplots and Table 2 gives
summary statistics. Discussion of the results is given in the text.

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References

- Alexe, M., Bergamaschi, P., Segers, A., Detmers, R., Butz, A., Hasekamp, O., Guerlet, S., Parker, R.,
 450 Boesch, H., Frankenberg, C., Scheepmaker, R. A., Dlugokencky, E., Sweeney, C., Wofsy, S. C., and
 Kort, E. A.: Inverse modelling of CH₄ emissions for 2010–2011 using different satellite retrieval products
 from GOSAT and SCIAMACHY, *Atmos. Chem. Phys.*, 15, 113–133, doi:10.5194/acp-15-113-2015, 2015.
- [Andrews, A. E., Kofler, J. D., Trudeau, M. E., Williams, J. C., Neff, D. H., Masarie, K. A., Chao, D. Y.,
 Kitzis, D. R., Novelli, P. C., Zhao, C. L., Dlugokencky, E. J., Lang, P. M., Crotwell, M. J., Fischer, M. L.,
 455 Parker, M. J., Lee, J. T., Baumann, D. D., Desai, A. R., Stanier, C. O., De Wekker, S. F. J., Wolfe, D. E.,
 Munger, J. W., and Tans, P. P.: CO₂, CO, and CH₄ measurements from tall towers in the NOAA Earth
 System Research Laboratory's Global Greenhouse Gas Reference Network: instrumentation, uncertainty
 analysis, and recommendations for future high-accuracy greenhouse gas monitoring efforts, *Atmospheric
 Measurement Techniques*, 7, 647–687, doi:10.5194/amt-7-647-2014, 2014.](#)
- 460 Bergamaschi, P., Frankenberg, C., Meirink, J. F., Krol, M., Dentener, F., Wagner, T., Platt, U., Kaplan, J. O.,
 Körner, S., Heimann, M., Dlugokencky, E. J., and Goede, A.: Satellite cartography of atmospheric methane
 from SCIAMACHY on board ENVISAT: 2. Evaluation based on inverse model simulations, *J. Geophys.
 Res.*, 112, D02304, doi:10.1029/2006jd007268, 2007.
- Bergamaschi, P., Frankenberg, C., Meirink, J. F., Krol, M., Villani, M. G., Houweling, S., Dentener, F.,
 465 Dlugokencky, E. J., Miller, J. B., Gatti, L. V., Engel, A., and Levin, I.: Inverse modeling of global
 and regional CH₄ emissions using SCIAMACHY satellite retrievals, *J. Geophys. Res.*, 114, D22301,
 doi:10.1029/2009jd012287, 2009.
- Bergamaschi, P., Houweling, S., Segers, A., Krol, M., Frankenberg, C., Scheepmaker, R. A., Dlugokencky, E.,
 Wofsy, S. C., Kort, E. A., Sweeney, C., Schuck, T., Brenninkmeijer, C., Chen, H., Beck, V., and Ger-
 470 big, C.: Atmospheric CH₄ in the first decade of the 21st century: iverse modeling analysis using SCIA-
 MACHY satellite retrievals and NOAA surface measurements, *J. Geophys. Res.-Atmos.*, 118, 7350–7369,
 doi:10.1002/jgrd.50480, 2013.
- Biraud, S. C., Torn, M. S., Smith, J. R., Sweeney, C., Riley, W. J., and Tans, P. P.: A multi-year record of airborne
 CO₂ observations in the US Southern Great Plains, *Atmos. Meas. Tech.*, 6, 751–763, doi:10.5194/amt-6-751-
 475 2013, 2013.
- Bishop, C. M.: *Pattern Recognition and Machine Learning*, 1st edn., Springer, New York, 2007.
- Bruhwyler, L., Dlugokencky, E., Masarie, K., Ishizawa, M., Andrews, A., Miller, J., Sweeney, C., Tans, P., and
 Worthy, D.: CarbonTracker-CH₄: an assimilation system for estimating emissions of atmospheric methane,
Atmos. Chem. Phys., 14, 8269–8293, doi:10.5194/acp-14-8269-2014, 2014.
- 480 Butz, A., Hasekamp, O. P., Frankenberg, C., Vidot, J., and Aben, I.: CH₄ retrievals from space-based solar
 backscatter measurements: performance evaluation against simulated aerosol and cirrus loaded scenes, *J.
 Geophys. Res.*, 115, D24302, doi:10.1029/2010jd014514, 2010.
- Butz, A., Guerlet, S., Hasekamp, O., Schepers, D., Galli, A., Aben, I., Frankenberg, C., Hartmann, J. M.,
 Tran, H., Kuze, A., Keppel-Aleks, G., Toon, G., Wunch, D., Wennberg, P., Deutscher, N., Griffith, D.,
 485 Macatangay, R., Messerschmidt, J., Notholt, J., and Warneke, T.: Toward accurate CO₂ and CH₄ obser-
 vations from GOSAT, *Geophys. Res. Lett.*, 38, L14812, doi:10.1029/2011gl047888, 2011.

- Caulton, D. R., Shepson, P. B., Santoro, R. L., Sparks, J. P., Howarth, R. W., Ingraffea, A. R., Cambaliza, M. O., Sweeney, C., Karion, A., Davis, K. J., Stirm, B. H., Montzka, S. A., and Miller, B. R.: Toward a better understanding and quantification of methane emissions from shale gas development, *P. Natl. Acad. Sci. USA*, 111, 6237–6242, doi:10.1073/pnas.1316546111, 2014.
- Climate and Clean Air Coalition: 2013–2014 Annual Report, Tech. rep., available at: http://ccacoalition.org/docs/pdf/CCAC_Annual_Report_2013-2014.pdf, 1 December 2014.
- Cressot, C., Chevallier, F., Bousquet, P., Crevoisier, C., Dlugokencky, E. J., Fortems-Cheiney, A., Frankenberg, C., Parker, R., Pison, I., Scheepmaker, R. A., Montzka, S. A., Krummel, P. B., Steele, L. P., and Langenfelds, R. L.: On the consistency between global and regional methane emissions inferred from SCIAMACHY, TANSO-FTS, IASI and surface measurements, *Atmos. Chem. Phys.*, 14, 577–592, doi:10.5194/acp-14-577-2014, 2014.
- Deutscher, N. M., Griffith, D. W. T., Bryant, G. W., Wennberg, P. O., Toon, G. C., Washenfelder, R. A., Keppel-Aleks, G., Wunch, D., Yavin, Y., Allen, N. T., Blavier, J.-F., Jiménez, R., Daube, B. C., Bright, A. V., Matross, D. M., Wofsy, S. C., and Park, S.: Total column CO₂ measurements at Darwin, Australia – site description and calibration against in situ aircraft profiles, *Atmos. Meas. Tech.*, 3, 947–958, doi:10.5194/amt-3-947-2010, 2010.
- Dlugokencky, E. J., Nisbet, E. G., Fisher, R., and Lowry, D.: Global atmospheric methane: budget, changes and dangers, *Philos. T. Roy. Soc. A*, 369, 2058–2072, doi:10.1098/rsta.2010.0341, 2011.
- EPA, U.: Inventory of U.S. Greenhouse Gas Emissions and Sinks: 1990–2012, Tech. rep., U.S. Environmental Protection Agency, Washington DC, 2014.
- European Commission: Emission Database for Global Atmospheric Research (EDGAR), release version 4.2, Tech. rep., Joint Research Centre (JRC)/Netherlands Environmental Assessment Agency (PBL), available at: <http://edgar.jrc.ec.europa.eu> (last access: 1 December 2014), 2011.
- Frankenberg, C., Meirink, J. F., van Weele, M., Platt, U., and Wagner, T.: Assessing methane emissions from global space-borne observations, *Science*, 308, 1010–1014, doi:10.1126/science.1106644, 2005.
- Frankenberg, C., Aben, I., Bergamaschi, P., Dlugokencky, E. J., van Hees, R., Houweling, S., van der Meer, P., Snel, R., and Tol, P.: Global column-averaged methane mixing ratios from 2003 to 2009 as derived from SCIAMACHY: trends and variability, *J. Geophys. Res.*, 116, D04302, doi:10.1029/2010jd014849, 2011.
- Fraser, A., Palmer, P. I., Feng, L., Boesch, H., Cogan, A., Parker, R., Dlugokencky, E. J., Fraser, P. J., Krummel, P. B., Langenfelds, R. L., O’Doherty, S., Prinn, R. G., Steele, L. P., van der Schoot, M., and Weiss, R. F.: Estimating regional methane surface fluxes: the relative importance of surface and GOSAT mole fraction measurements, *Atmos. Chem. Phys.*, 13, 5697–5713, doi:10.5194/acp-13-5697-2013, 2013.
- Fung, I., John, J., Lerner, J., Matthews, E., Prather, M., Steele, L. P., and Fraser, P. J.: Three-dimensional model synthesis of the global methane cycle, *J. Geophys. Res.*, 96, 13033, doi:10.1029/91jd01247, 1991.
- Geibel, M. C., Messerschmidt, J., Gerbig, C., Blumenstock, T., Chen, H., Hase, F., Kolle, O., Lavrič, J. V., Notholt, J., Palm, M., Rettinger, M., Schmidt, M., Sussmann, R., Warneke, T., and Feist, D. G.: Calibration of column-averaged CH₄ over European TCCON FTS sites with airborne in-situ measurements, *Atmos. Chem. Phys.*, 12, 8763–8775, doi:10.5194/acp-12-8763-2012, 2012.
- Heald, C. L., Jacob, D. J., Jones, D. B. A., Palmer, P. I., Logan, J. A., Streets, D. G., Sachse, G. W., Gille, J. C., Hoffman, R. N., and Nehrkorn, T.: Comparative inverse analysis of satellite (MOPITT) and aircraft (TRACE-

P) observations to estimate Asian sources of carbon monoxide, *J. Geophys. Res.-Atmos.*, 109, D23306, doi:10.1029/2004jd005185, 2004.

Henze, D. K., Hakami, A., and Seinfeld, J. H.: Development of the adjoint of GEOS-Chem, *Atmos. Chem. Phys.*, 7, 2413–2433, doi:10.5194/acp-7-2413-2007, 2007.

Hsu, Y. K., VanCuren, T., Park, S., Jakober, C., Herner, J., FitzGibbon, M., Blake, D. R., and Parrish, D. D.: Methane emissions inventory verification in southern California, *Atmos. Environ.*, 44, 1–7, doi:10.1016/J.Atmosenv.2009.10.002, 2010.

Hudman, R. C., Jacob, D. J., Turquety, S., Leibensperger, E. M., Murray, L. T., Wu, S., Gilliland, A. B., Avery, M., Bertram, T. H., Brune, W., Cohen, R. C., Dibb, J. E., Flocke, F. M., Fried, A., Holloway, J., Neuman, J. A., Orville, R., Perring, A., Ren, X., Sachse, G. W., Singh, H. B., Swanson, A., and Wooldridge, P. J.: Surface and lightning sources of nitrogen oxides over the United States: magnitudes, chemical evolution, and outflow, *J. Geophys. Res.*, 112, D12S05, doi:10.1029/2006jd007912, 2007.

IPCC: Climate Change 2013: The Physical Science Basis. Contribution of Working Group I to the Fifth Assessment Report of the Intergovernmental Panel on Climate Change, Tech. rep., Cambridge, United Kingdom and New York, 2013.

Jeong, S., Zhao, C., Andrews, A. E., Bianco, L., Wilczak, J. M., and Fischer, M. L.: Seasonal variation of CH₄ emissions from central California, *J. Geophys. Res.*, 117, D11306, doi:10.1029/2011jd016896, 2012.

Jeong, S., Hsu, Y.-K., Andrews, A. E., Bianco, L., Vaca, P., Wilczak, J. M., and Fischer, M. L.: A multitower measurement network estimate of California's methane emissions, *J. Geophys. Res.-Atmos.*, 118, 11339–11351, doi:10.1002/jgrd.50854, 2013.

Kaplan, J. O.: Wetlands at the last glacial maximum: distribution and methane emissions, *Geophys. Res. Lett.*, 29, 3-1–3-4, doi:10.1029/2001gl013366, 2002.

Karion, A., Sweeney, C., Pétron, G., Frost, G., Michael Hardesty, R., Kofler, J., Miller, B. R., Newberger, T., Wolter, S., Banta, R., Brewer, A., Dlugokencky, E., Lang, P., Montzka, S. A., Schnell, R., Tans, P., Trainer, M., Zamora, R., and Conley, S.: Methane emissions estimate from airborne measurements over a western United States natural gas field, *Geophys. Res. Lett.*, 40, 4393–4397, doi:10.1002/grl.50811, 2013.

Katzenstein, A. S., Doezeema, L. A., Simpson, I. J., Blake, D. R., and Rowland, F. S.: Extensive regional atmospheric hydrocarbon pollution in the southwestern United States, *P. Natl. Acad. Sci. USA*, 100, 11975–11979, doi:10.1073/pnas.1635258100, 2003.

Kirschke, S., Bousquet, P., Ciais, P., Saunio, M., Canadell, J. G., Dlugokencky, E. J., Bergamaschi, P., Bergmann, D., Blake, D. R., Bruhwiler, L., Cameron-Smith, P., Castaldi, S., Chevallier, F., Feng, L., Fraser, A., Heimann, M., Hodson, E. L., Houweling, S., Josse, B., Fraser, P. J., Krummel, P. B., Lamarque, J.-F., Langenfelds, R. L., Le Quéré, C., Naik, V., O'Doherty, S., Palmer, P. I., Pison, I., Plummer, D., Poulter, B., Prinn, R. G., Rigby, M., Ringeval, B., Santini, M., Schmidt, M., Shindell, D. T., Simpson, I. J., Spahni, R., Steele, L. P., Strode, S. A., Sudo, K., Szopa, S., van der Werf, G. R., Voulgarakis, A., van Weele, M., Weiss, R. F., Williams, J. E., and Zeng, G.: Three decades of global methane sources and sinks, *Nat. Geosci.*, 6, 813–823, doi:10.1038/ngeo1955, 2013.

Kort, E. A., Eluszkiewicz, J., Stephens, B. B., Miller, J. B., Gerbig, C., Nehrkorn, T., Daube, B. C., Kaplan, J. O., Houweling, S., and Wofsy, S. C.: Emissions of CH₄ and N₂O over the United States and Canada based on

a receptor-oriented modeling framework and COBRA-NA atmospheric observations, *Geophys. Res. Lett.*, 35, L18808, doi:10.1029/2008gl034031, 2008.

Kort, E. A., Frankenberg, C., Costigan, K. R., Lindenmaier, R., Dubey, M. K., and Wunch, D.: Four corners: the largest US methane anomaly viewed from space, *Geophys. Res. Lett.*, 41, 6898–6903, doi:10.1002/2014gl061503, 2014.

Kuze, A., Suto, H., Nakajima, M., and Hamazaki, T.: Thermal and near infrared sensor for carbon observation Fourier-transform spectrometer on the Greenhouse Gases Observing Satellite for greenhouse gases monitoring, *Appl. Optics*, 48, 6716–6733, doi:10.1364/AO.48.006716, 2009.

Messerschmidt, J., Geibel, M. C., Blumenstock, T., Chen, H., Deutscher, N. M., Engel, A., Feist, D. G., Gerbig, C., Gisi, M., Hase, F., Katrynski, K., Kolle, O., Lavrič, J. V., Notholt, J., Palm, M., Ramonet, M., Rettinger, M., Schmidt, M., Sussmann, R., Toon, G. C., Truong, F., Warneke, T., Wennberg, P. O., Wunch, D., and Xueref-Remy, I.: Calibration of TCCON column-averaged CO₂: the first aircraft campaign over European TCCON sites, *Atmos. Chem. Phys.*, 11, 10765–10777, doi:10.5194/acp-11-10765-2011, 2011.

Messerschmidt, J., Chen, H., Deutscher, N. M., Gerbig, C., Grupe, P., Katrynski, K., Koch, F.-T., Lavrič, J. V., Notholt, J., Rödenbeck, C., Ruhe, W., Warneke, T., and Weinzierl, C.: Automated ground-based remote sensing measurements of greenhouse gases at the Biaystok site in comparison with collocated in situ measurements and model data, *Atmos. Chem. Phys.*, 12, 6741–6755, doi:10.5194/acp-12-6741-2012, 2012.

Miller, S. M., Wofsy, S. C., Michalak, A. M., Kort, E. A., Andrews, A. E., Biraud, S. C., Dlugokencky, E. J., Eluszkiewicz, J., Fischer, M. L., Janssens-Maenhout, G., Miller, B. R., Miller, J. B., Montzka, S. A., Nehrkorn, T., and Sweeney, C.: Anthropogenic emissions of methane in the United States, *P. Natl. Acad. Sci. USA*, 110, 20018–20022, doi:10.1073/pnas.1314392110, 2013.

Monteil, G., Houweling, S., Butz, A., Guerlet, S., Schepers, D., Hasekamp, O., Frankenberg, C., Scheepmaker, R., Aben, I., and Röckmann, T.: Comparison of CH₄ inversions based on 15 months of GOSAT and SCIAMACHY observations, *J. Geophys. Res.-Atmos.*, 118, 11807–11823, doi:10.1002/2013jd019760, 2013.

Park, R. J., Jacob, D. J., Field, B. D., Yantosca, R. M., and Chin, M.: Natural and transboundary pollution influences on sulfate-nitrate-ammonium aerosols in the United States: implications for policy, *J. Geophys. Res.-Atmos.*, 109, D15204, doi:10.1029/2003jd004473, 2004.

Parker, R., Boesch, H., Cogan, A., Fraser, A., Feng, L., Palmer, P. I., Messerschmidt, J., Deutscher, N., Griffith, D. W. T., Notholt, J., Wennberg, P. O., and Wunch, D.: Methane observations from the Greenhouse Gases Observing SATellite: comparison to ground-based TCCON data and model calculations, *Geophys. Res. Lett.*, 38, L15807, doi:10.1029/2011gl047871, 2011.

Peischl, J., Ryerson, T. B., Holloway, J. S., Trainer, M., Andrews, A. E., Atlas, E. L., Blake, D. R., Daube, B. C., Dlugokencky, E. J., Fischer, M. L., Goldstein, A. H., Guha, A., Karl, T., Kofler, J., Kosciuch, E., Misztal, P. K., Perring, A. E., Pollack, I. B., Santoni, G. W., Schwarz, J. P., Spackman, J. R., Wofsy, S. C., and Parrish, D. D.: Airborne observations of methane emissions from rice cultivation in the Sacramento Valley of California, *J. Geophys. Res.-Atmos.*, 117, D00V25, doi:10.1029/2012jd017994, 2012.

Peischl, J., Ryerson, T. B., Brioude, J., Aikin, K. C., Andrews, A. E., Atlas, E., Blake, D., Daube, B. C., de Gouw, J. A., Dlugokencky, E., Frost, G. J., Gentner, D. R., Gilman, J. B., Goldstein, A. H., Harley, R. A., Holloway, J. S., Kofler, J., Kuster, W. C., Lang, P. M., Novelli, P. C., Santoni, G. W., Trainer, M., Wofsy, S. C.,

and Parrish, D. D.: Quantifying sources of methane using light alkanes in the Los Angeles basin, California, *J. Geophys. Res.-Atmos.*, 118, 4974–4990, doi:10.1002/jgrd.50413, 2013.

Pickett-Heaps, C. A., Jacob, D. J., Wecht, K. J., Kort, E. A., Wofsy, S. C., Diskin, G. S., Worthy, D. E. J., Kaplan, J. O., Bey, I., and Drevet, J.: Magnitude and seasonality of wetland methane emissions from the Hudson Bay Lowlands (Canada), *Atmos. Chem. Phys.*, 11, 3773–3779, doi:10.5194/acp-11-3773-2011, 2011.

Prather, M. J., Holmes, C. D., and Hsu, J.: Reactive greenhouse gas scenarios: systematic exploration of uncertainties and the role of atmospheric chemistry, *Geophys. Res. Lett.*, 39, L09803, doi:10.1029/2012gl051440, 2012.

President’s Climate Action Plan: Tech. rep., The White House, available at: <http://www.whitehouse.gov/sites/default/files/image/president27sclimateactionplan.pdf> (last access: 1 December 2014), 2013.

President’s Climate Action Plan: Strategy to Reduce Methane Emissions, Tech. rep., The White House, available at: http://www.whitehouse.gov/sites/default/files/strategy_to_reduce_methane_emissions_2014-03-28_final.pdf, last access: 1 December 2014.

Ramanathan, V. and Xu, Y.: The Copenhagen Accord for limiting global warming: criteria, constraints, and available avenues, *P. Natl. Acad. Sci. USA*, 107, 8055–8062, doi:10.1073/pnas.1002293107, 2010.

Rodgers, C. D.: *Inverse Methods for Atmospheric Sounding*, World Scientific, Singapore, 2000.

[Saad, K. M., Wunch, D., Toon, G. C., Bernath, P., Boone, C., Coon, B., Deutscher, N. M., Griffith, D. W. T., Kivi, R., Notholt, J., Roehl, C., Schneider, M., Sherlock, V., and Wennberg, P. O.: Derivation of tropospheric methane from TCCON CH₄ and HF total column observations, *Atmos. Meas. Tech.*, 7, 2907–2918, doi:10.5194/amt-7-2907-2014, 2014.](#)

Santoni, G. W., Xiang, B., Kort, E. A., Daubel, B. C., Andrews, A. E., Sweeney, C., Wecht, K. J., Peischl, J., Ryerson, T. B., Angevine, W. M., Trainer, M., Nehrkorn, T., Eluszkiewicz, J., Jeong, S., Fischer, M. L., Ferrare, R. A., and Wofsy, S. C.: California’s methane budget derived from CalNex P-3 Aircraft Observations and a Lagrangian transport model, *J. Geophys. Res.*, submitted, 2014.

Schepers, D., Guerlet, S., Butz, A., Landgraf, J., Frankenberg, C., Hasekamp, O., Blavier, J. F., Deutscher, N. M., Griffith, D. W. T., Hase, F., Kyro, E., Morino, I., Sherlock, V., Sussmann, R., and Aben, I.: Methane retrievals from Greenhouse Gases Observing Satellite (GOSAT) shortwave infrared measurements: performance comparison of proxy and physics retrieval algorithms, *J. Geophys. Res.*, 117, D10307, doi:10.1029/2012jd017549, 2012.

Shindell, D., Kuylensstierna, J. C., Vignati, E., van Dingenen, R., Amann, M., Klimont, Z., Anenberg, S. C., Muller, N., Janssens-Maenhout, G., Raes, F., Schwartz, J., Faluvegi, G., Pozzoli, L., Kupiainen, K., Hoglund-Isaksson, L., Emberson, L., Streets, D., Ramanathan, V., Hicks, K., Oanh, N. T., Milly, G., Williams, M., Demkine, V., and Fowler, D.: Simultaneously mitigating near-term climate change and improving human health and food security, *Science*, 335, 183–189, doi:10.1126/science.1210026, 2012.

Turner, A. J. and Jacob, D. J.: Balancing aggregation and smoothing errors in inverse models, *Atmos. Chem. Phys. Discuss.*, 15, 1001–1026, doi:10.5194/acpd-15-1001-2015, 2015.

van der Werf, G. R., Randerson, J. T., Giglio, L., Collatz, G. J., Mu, M., Kasibhatla, P. S., Morton, D. C., DeFries, R. S., Jin, Y., and van Leeuwen, T. T.: Global fire emissions and the contribution of deforestation, savanna, forest, agricultural, and peat fires (1997–2009), *Atmos. Chem. Phys.*, 10, 11707–11735, doi:10.5194/acp-10-11707-2010, 2010.

[Wang, Z., Deutscher, N. M., Warneke, T., Notholt, J., Dils, B., Griffith, D. W. T., Schmidt, M., Ramonet, M., and Gerbig, C.: Retrieval of tropospheric column-averaged CH₄ mole fraction by solar absorption FTIR-spectrometry using N₂O as a proxy, *Atmospheric Measurement Techniques*, 7, 3295–3305, doi:10.5194/amt-7-3295-2014, 2014.](#)

- 650 Washenfelder, R. A., Toon, G. C., Blavier, J. F., Yang, Z., Allen, N. T., Wennberg, P. O., Vay, S. A., Matross, D. M., and Daube, B. C.: Carbon dioxide column abundances at the Wisconsin Tall Tower site, *J. Geophys. Res.*, 111, D22305, doi:10.1029/2006jd007154, 2006.
- Wecht, K. J., Jacob, D. J., Wofsy, S. C., Kort, E. A., Worden, J. R., Kulawik, S. S., Henze, D. K., Kopacz, M., and Payne, V. H.: Validation of TES methane with HIPPO aircraft observations: implications for inverse
655 modeling of methane sources, *Atmos. Chem. Phys.*, 12, 1823–1832, doi:10.5194/acp-12-1823-2012, 2012.
- Wecht, K. J., Jacob, D. J., Frankenberg, C., Jiang, Z., and Blake, D. R.: Mapping of North American methane emissions with high spatial resolution by inversion of SCIAMACHY satellite data, *J. Geophys. Res.-Atmos.*, 119, 7741–7756, doi:10.1002/2014jd021551, 2014a.
- Wecht, K. J., Jacob, D. J., Sulprizio, M. P., Santoni, G. W., Wofsy, S. C., Parker, R., Bösch, H., and Worden, J.:
660 Spatially resolving methane emissions in California: constraints from the CalNex aircraft campaign and from present (GOSAT, TES) and future (TROPOMI, geostationary) satellite observations, *Atmos. Chem. Phys.*, 14, 8173–8184, doi:10.5194/acp-14-8173-2014, 2014b.
- Wennberg, P. O., Mui, W., Wunch, D., Kort, E. A., Blake, D. R., Atlas, E. L., Santoni, G. W., Wofsy, S. C., Diskin, G. S., Jeong, S., and Fischer, M. L.: On the sources of methane to the Los Angeles atmosphere,
665 *Environ. Sci. Technol.*, 46, 9282–9289, doi:10.1021/es301138y, 2012.
- Wofsy, S. C.: HIPER Pole-to-Pole Observations (HIPPO): fine-grained, global-scale measurements of climatically important atmospheric gases and aerosols, *Philos. T. Roy. Soc. A*, 369, 2073–2086, doi:10.1098/rsta.2010.0313, 2011.
- Wunch, D., Wennberg, P. O., Toon, G. C., Keppel-Aleks, G., and Yavin, Y. G.: Emissions of greenhouse gases
670 from a North American megacity, *Geophys. Res. Lett.*, 36, L15810, doi:10.1029/2009gl039825, 2009.
- Wunch, D., Toon, G. C., Wennberg, P. O., Wofsy, S. C., Stephens, B. B., Fischer, M. L., Uchino, O., Abshire, J. B., Bernath, P., Biraud, S. C., Blavier, J.-F. L., Boone, C., Bowman, K. P., Browell, E. V., Campos, T., Connor, B. J., Daube, B. C., Deutscher, N. M., Diao, M., Elkins, J. W., Gerbig, C., Gottlieb, E., Griffith, D. W. T., Hurst, D. F., Jiménez, R., Keppel-Aleks, G., Kort, E. A., Macatangay, R., Machida, T., Matsueda, H., Moore, F., Morino, I., Park, S., Robinson, J., Roehl, C. M., Sawa, Y., Sherlock, V., Sweeney, C.,
675 Tanaka, T., and Zondlo, M. A.: Calibration of the Total Carbon Column Observing Network using aircraft profile data, *Atmos. Meas. Tech.*, 3, 1351–1362, doi:10.5194/amt-3-1351-2010, 2010.
- Wunch, D., Toon, G. C., Blavier, J. F., Washenfelder, R. A., Notholt, J., Connor, B. J., Griffith, D. W., Sherlock, V., and Wennberg, P. O.: The total carbon column observing network, *Philos. T. Roy. Soc. A*, 369,
680 2087–112, doi:10.1098/rsta.2010.0240, 2011.
- Xiao, Y., Logan, J. A., Jacob, D. J., Hudman, R. C., Yantosca, R., and Blake, D. R.: Global budget of ethane and regional constraints on US sources, *J. Geophys. Res.*, 113, D21306, doi:10.1029/2007jd009415, 2008.
- Yevich, R. and Logan, J. A.: An assessment of biofuel use and burning of agricultural waste in the developing world, *Global Biogeochem. Cy.*, 17, 1095, doi:10.1029/2002gb001952, 2003.

Table 1. 2009–2011 methane emissions^a.

Source Type	Contiguous US		North America		Global	
	Prior	Posterior ^b	Prior	Posterior ^b	Prior	Posterior
Total	31.4	51.3–52.5	63.3	88.5–91.3	537	539
Wetlands	5.9	9.0–10.1	20.4	22.9–23.7	164	169
Livestock	8.9	12.6–17.0	14.5	20.0–25.7	111	116
Oil/Gas	5.4	8.7–13.4	10.8	15.5–22.3	69	67
Waste ^c	5.5	8.0–8.5	9.7	13.4–14.5	60	65
Coal	4.0	4.7–6.5	4.3	5.0–6.8	47	30
Rice	0.4	0.8–0.9	0.5	0.9–1.0	38	45
Open Fires	0.1	0.1	1.0	0.9	17	16
Other ^d	1.1	1.6–1.7	2.2	3.0–3.3	31	32
Natural ^e	7.5	9.8–11.1	25.0	25.1–26.2	176	181
Anthropogenic ^f	25.0	40.2–42.7	41.9	62.3–66.2	361	358

^aEmissions are in Tg a^{-1} . Prior emissions are mainly from EDGARv4.2 for anthropogenic sources and Pickett-Heaps et al. (2011) for wetlands (see Appendix).

^bRange from two inversions with different assumptions for prior error (see text).

^cIncluding landfills and waste water.

^dIncluding fuel combustion, termites, and soil absorption.

^eIncluding wetlands, open fires, termites, and soil absorption.

^fIncluding livestock, oil/gas, waste, coal, rice, and fuel combustion.

Table 2. GEOS-Chem comparison to 2009–2011 suborbital methane observations^a.

Observations	R^2		Slope ^b		Mean Bias ^c	
	Prior	Posterior	Prior	Posterior	Prior	Posterior
Global						
HIPPO (II-V)	0.94	0.94	0.97	0.95	1.5 <u>4.2</u>	2.0 <u>4.4</u>
TCCON	0.77 <u>0.82</u>	0.77 <u>0.83</u>	0.93 <u>0.94</u>	0.98	3.8 <u>6.4</u>	7.0 <u>8.1</u>
NOAA Cooperative Flask Network <u>NOAA/ESRL Surface Flasks</u>	0.62 <u>0.66</u>	0.65 <u>0.66</u>	0.98 <u>1.08</u>	0.95 <u>1.04</u>	6.7 <u>16.1</u>	7.9 <u>14.1</u>
North America						
NOAA Cooperative Flask <u>NOAA/ESRL Tall Tower Network</u>	0.57 <u>0.40</u>	0.60 <u>0.48</u>	0.64 <u>0.72</u>	1.00 <u>1.03</u>	-1.3 <u>13.3</u>	9.6 <u>3.1</u>
NOAA/ DOE <u>ESRL</u> Aircraft Program	0.56 <u>0.54</u>	0.64 <u>0.61</u>	0.71 <u>0.75</u>	0.92 <u>0.94</u>	-0.7 <u>0.2</u>	6.5 <u>6.7</u>
<u>NOAA/ESRL Surface Flasks</u>	<u>0.60</u>	<u>0.67</u>	<u>0.67</u>	<u>1.01</u>	<u>-5.6</u>	<u>7.1</u>

^aGEOS-Chem at $4^\circ \times 5^\circ$ resolution globally and $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution for North America, using either prior emissions (Table 1 and Figs. A3 and A4) or posterior emissions optimized with the inversion. Further details on the comparisons are in Figs. A3–A5. NOAA observations are from the NOAA/ESRL Greenhouse Gas Reference Network. References for the observations are given in the text.

^bSlope (in ppbv ppbv⁻¹) is from a reduced-major-axis (RMA) regression.

^cMean bias is the mean difference (in ppbv) between model and observations.

- 685 Zhao, C. F., Andrews, A. E., Bianco, L., Eluszkiewicz, J., Hirsch, A., MacDonald, C., Nehrkorn, T., and Fischer, M. L.: Atmospheric inverse estimates of methane emissions from Central California, J. Geophys. Res.-Atmos., 114,D16302, doi:10.1029/2008jd011671, 2009.

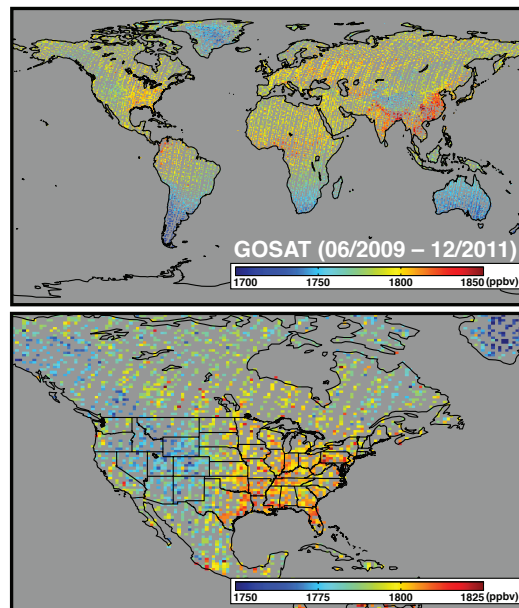


Figure 1. Mean GOSAT observations of the weighted column-average methane dry-mole fraction (X_{CH_4}) for June 2009–December 2011, globally and for North America. The data are version 4 proxy methane retrievals from Parker et al. (2011) that pass all quality flags (<http://www.leos.le.ac.uk/GHG/data/styled/index.html>).

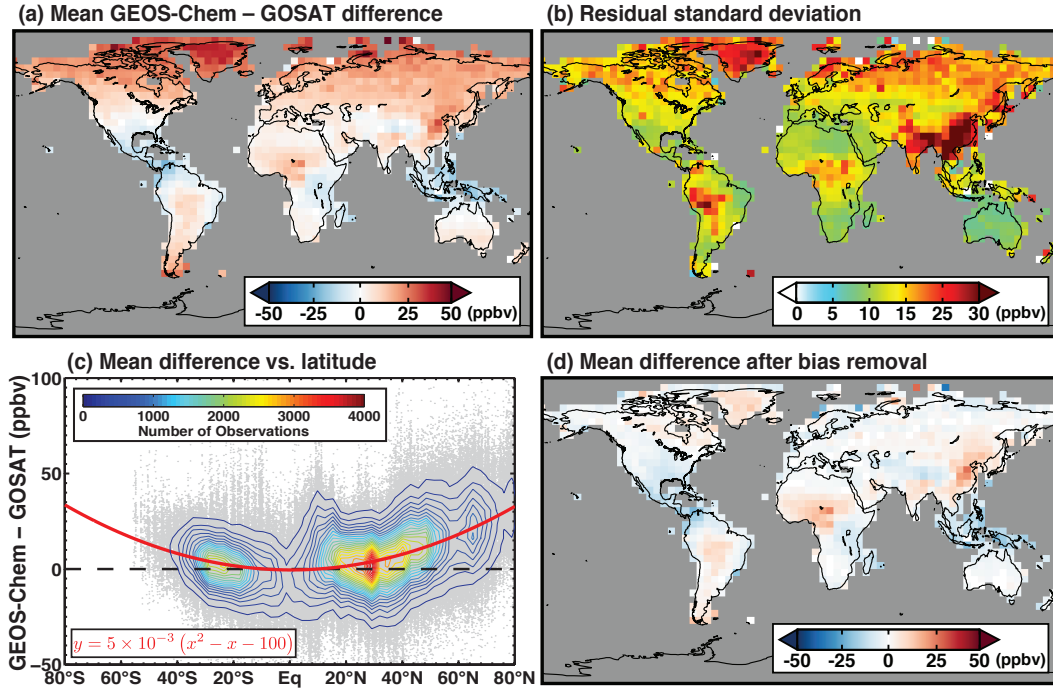


Figure 2. Comparison of the GOSAT observations from Fig. 1 to the GEOS-Chem model with prior emissions. The top panels show the mean bias and residual SD for the model-satellite difference. The bottom left panel shows the model-satellite difference as a function of latitude for individual observations along with the data density (contours), and a quadratic regression (red line; x in degrees latitude, y in ppbv) as estimate of the bias. The regression excludes grid squares with residual SD in excess of 20 ppbv as model bias in prior emissions could dominate the difference. The bottom right panel is the same as top left but using the bias correction from the bottom left panel.

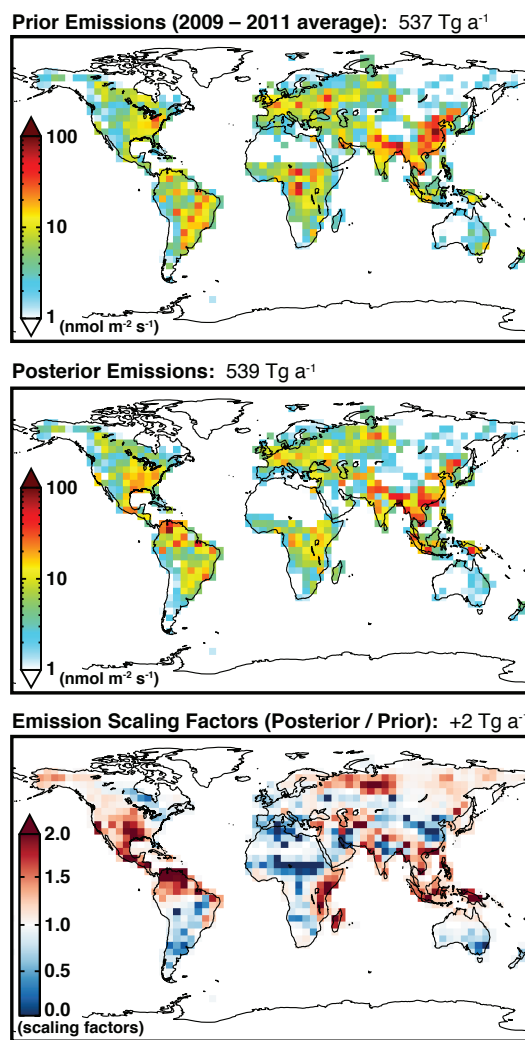


Figure 3. Optimization of methane emissions for 2009–2011 at $4^\circ \times 5^\circ$ horizontal resolution using GOSAT observations. The panels show prior emissions, posterior emissions, and the ratio between the two.

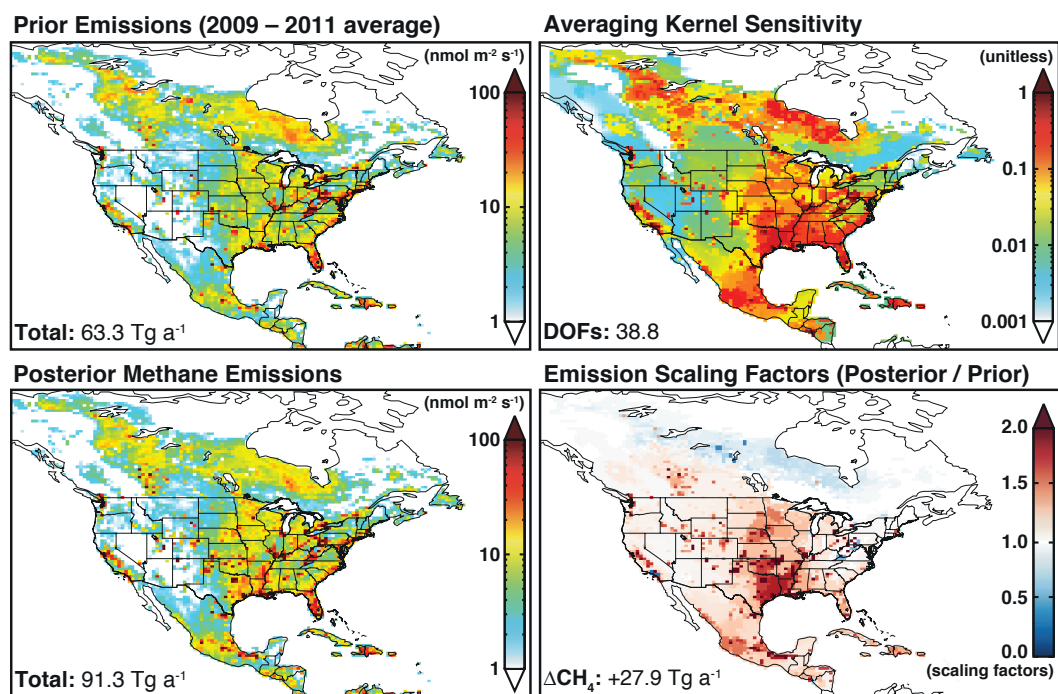


Figure 4. Methane emissions in North America in 2009–2011. The left panels show the prior and posterior emissions and the bottom right panel shows the scaling factors. The top right panel shows the diagonal elements of the averaging kernel matrix for the methane emission inversion. The degrees of freedom for signal (DOFS) is the trace of the averaging kernel matrix.

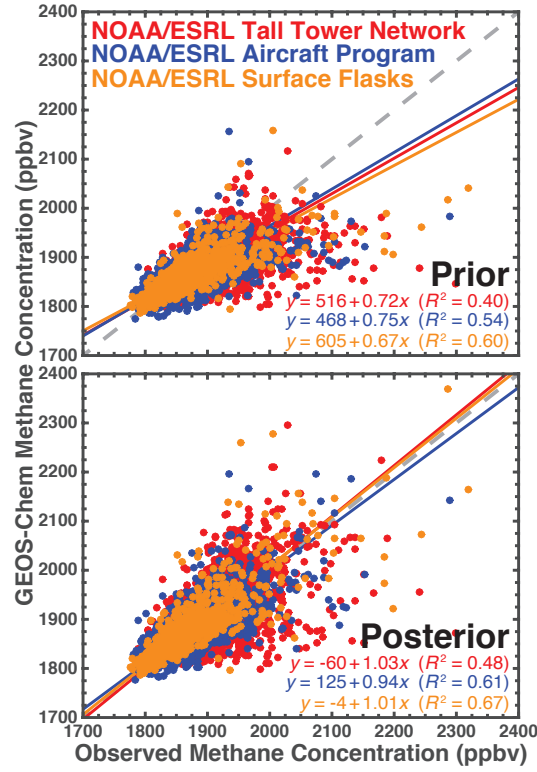


Figure 5. Evaluation of the GOSAT inversion of methane emissions for North America with independent data sets. The scatterplots show comparisons of GEOS-Chem ($\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution) methane concentrations with observations from the NOAA ~~cooperative surface flask~~ /ESRL tall tower network (red), NOAA/ESRL aircraft program (blue circles), and the NOAA/DOE vertical aircraft profiles-ESRL surface flask network (red squares/orange), using prior emissions (top) and posterior emissions (bottom). The 1 : 1 lines (dashed) and reduced-major-axis (RMA, solid) linear regressions are also shown. RMA regression parameters are shown inset and correspond to the statistics of Table 2.

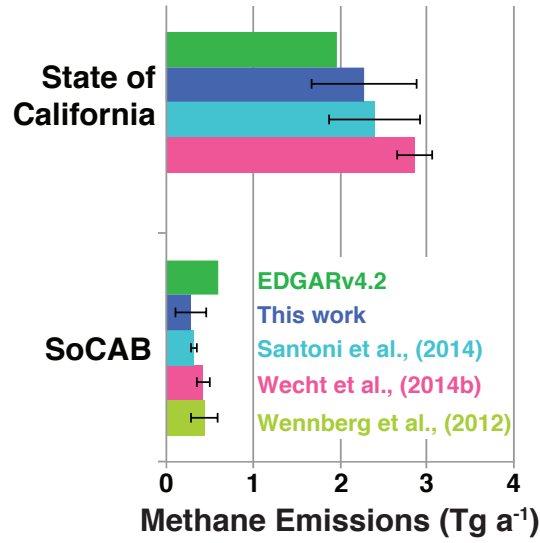


Figure 6. Methane emissions for the state of California (top) and for the Southern California Air Basin (SoCAB; bottom). Our posterior emissions (“this work”) are compared to prior emissions (EDGARv4.2) and to previous inverse estimates constrained by surface and aircraft observations. SoCAB is defined following Wennberg et al. (2012) as the domain 33.5°–34.5° N, 117°–119° W.

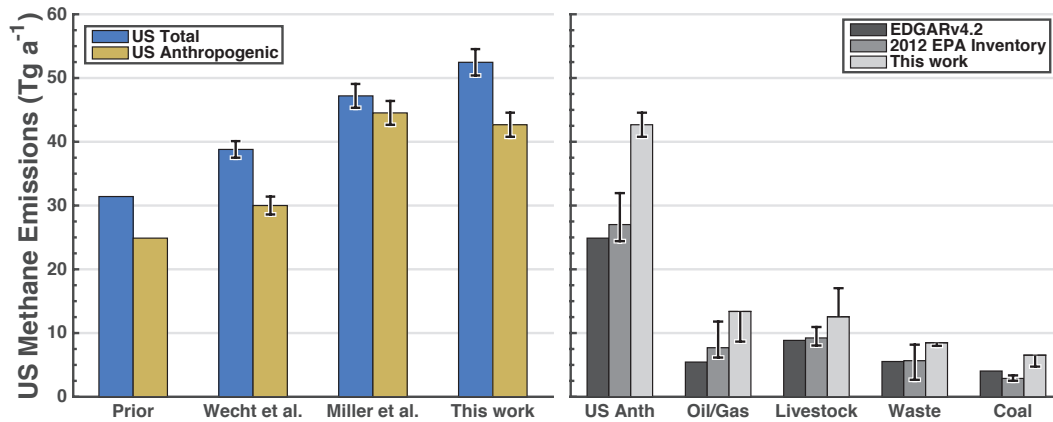


Figure 7. Methane emissions in the contiguous US. The left panel shows our best estimates of total and anthropogenic emissions (“this work”) compared to the prior (EDGARv4.2 for anthropogenic sources, Pickett-Heaps et al. (2011) for wetlands) and the previous inverse studies of Wecht et al. (2014a) and Miller et al. (2013). The right panel partitions US anthropogenic emissions by source types and compares our results (“this work”) to EDGARv4.2 and to the 2012 EPA inventory (EPA, 2014). Error bars on [sectoral emissions for](#) our results are defined by the sensitivity inversion with 30 % prior uncertainty for all state vector elements.

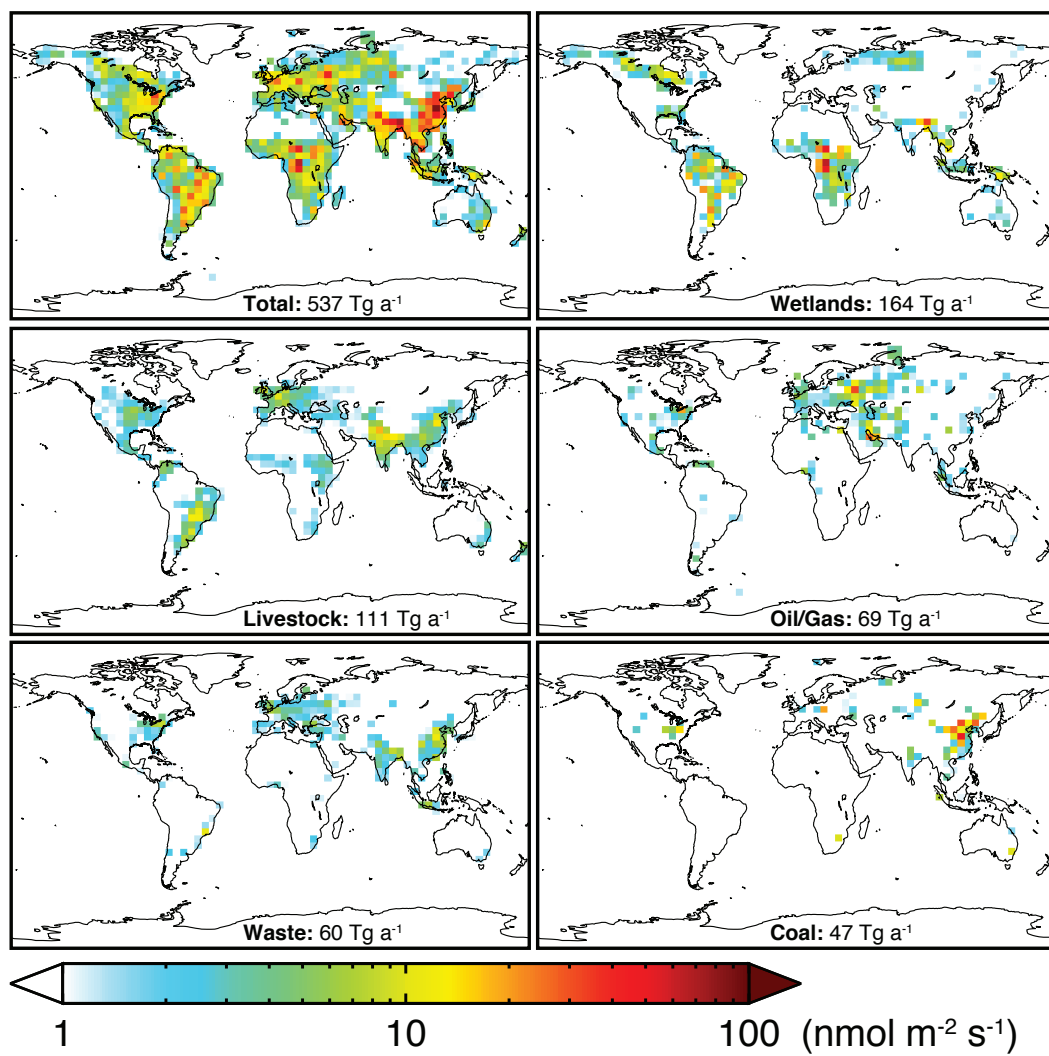


Figure A1. Prior 2009–2011 methane emissions used in the GEOS-Chem global simulation at $4^\circ \times 5^\circ$ resolution, and contributions from the top 5 sources.

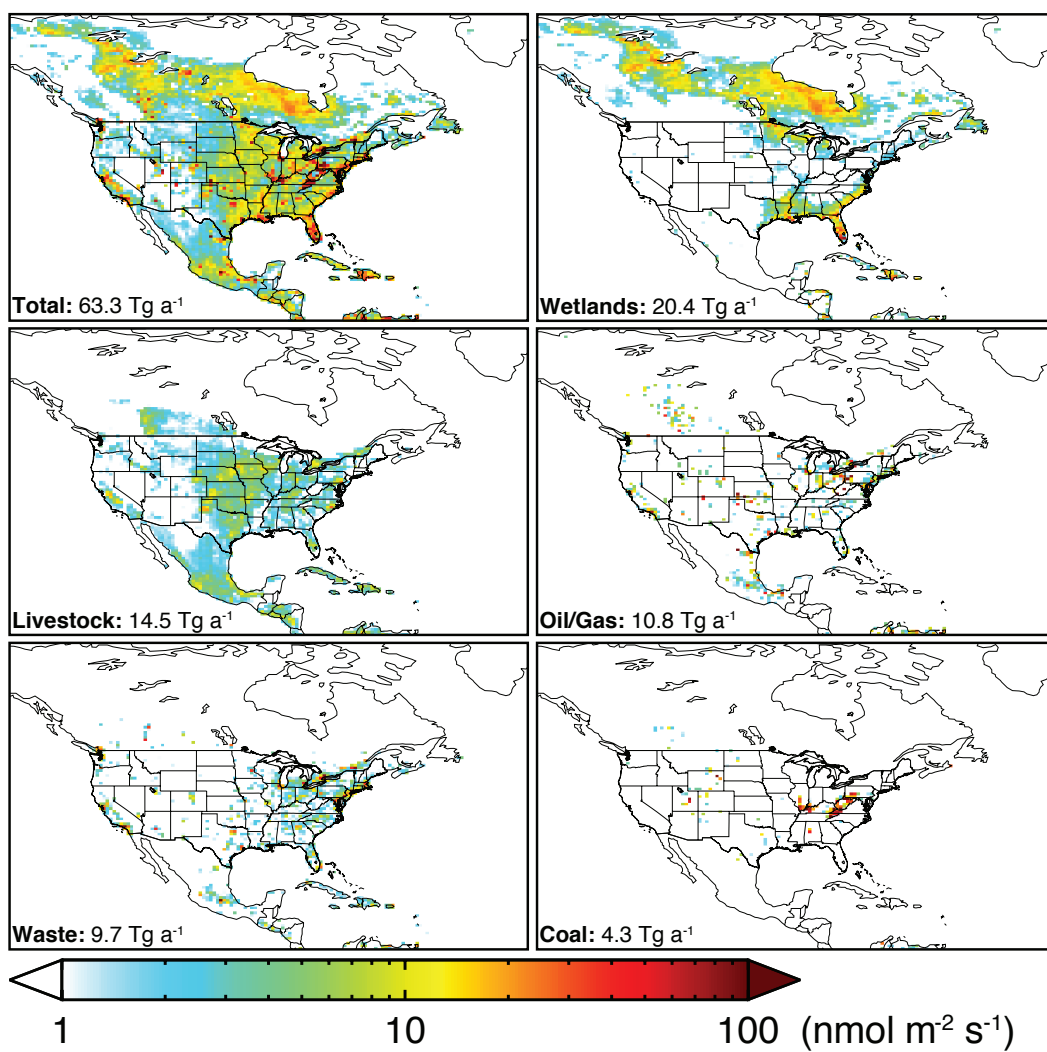


Figure A2. Same as Fig. A1 but for North America with $\frac{1}{2}^\circ \times \frac{2}{3}^\circ$ resolution.

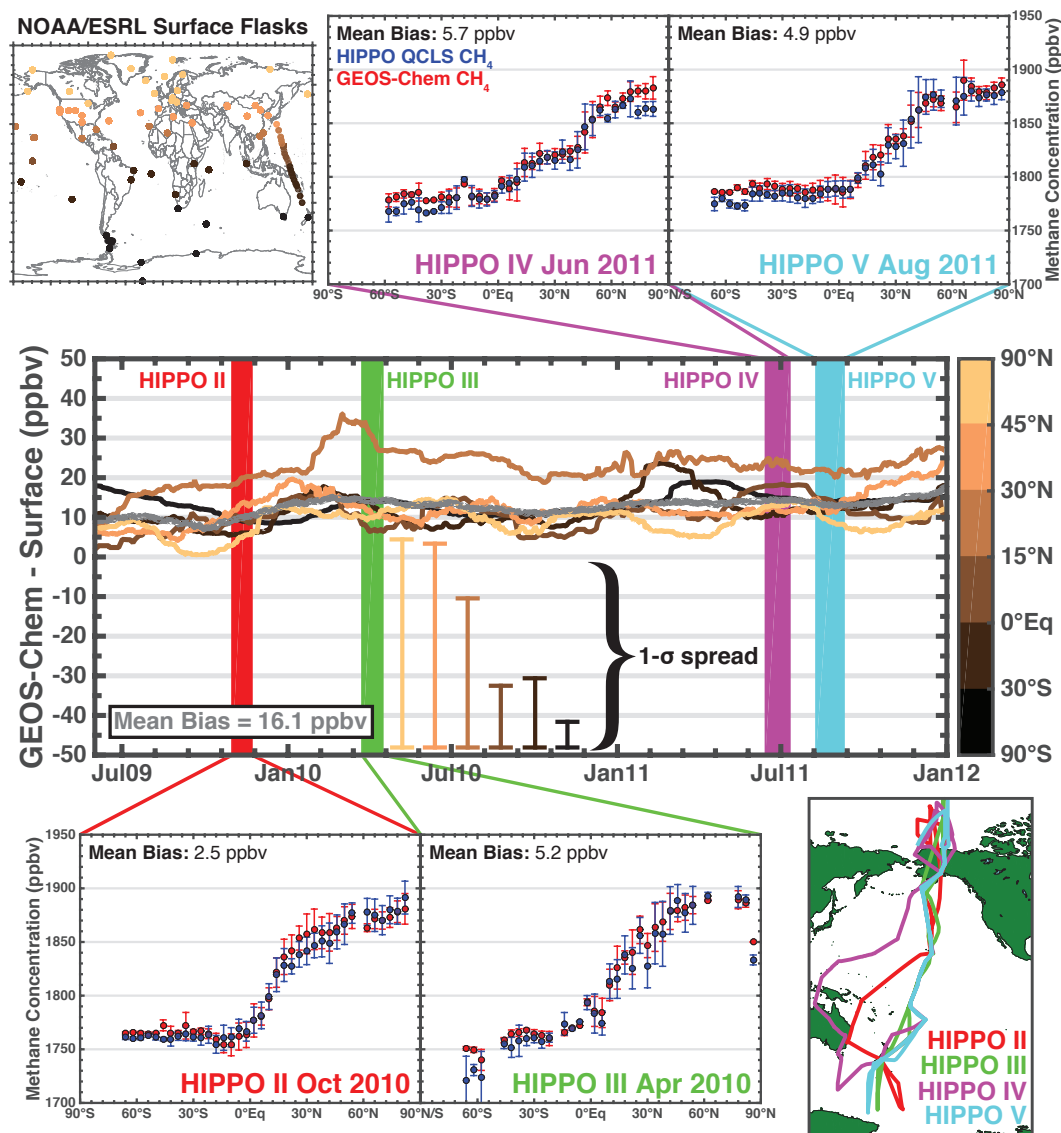


Figure A3. Global evaluation of the GEOS-Chem methane simulation at $4^\circ \times 5^\circ$ resolution (using prior emissions) with observations from the NOAA ~~cooperative~~-ESRL surface flask network (top left panel colored by latitude) and HIPPO aircraft deployments. The central panel shows 3-month running medians for 2009–2011 of the difference between GEOS-Chem and the flask data in different latitudinal bands. The gray line is for all of the observations. Error bars for the running medians are offset from the lines for clarity. Latitudinal profiles across the Pacific for the four HIPPO deployments over the period are also shown: in those panels the symbols represent the pressure-weighted tropospheric average concentrations and the vertical bars are the SD. Stratospheric air is excluded based on an ozone–CO concentration ratio larger than 1.25 (Hudman et al., 2007). Bottom right panel shows the HIPPO flight tracks.

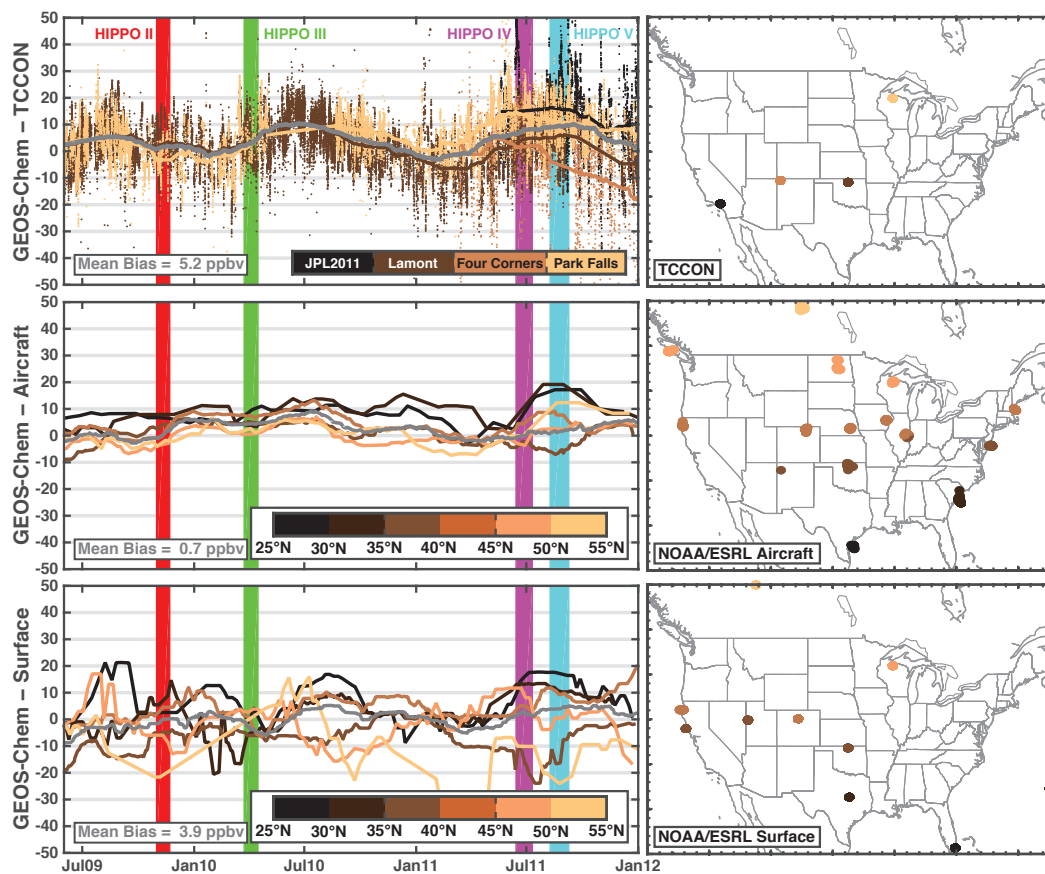


Figure A4. North American evaluation of the GEOS-Chem methane simulation at $4^\circ \times 5^\circ$ resolution (using prior emissions) with TCCON (top), NOAA/DOE-ESRL aircraft program (middle), and NOAAcooperative/ESRL surface flask network (bottom) observations over North America. Left panels show 3-month running medians for 2009–2011 of the difference between GEOS-Chem and the observations in different latitudinal bands and for all the data (gray line). Right panels show the location of the observations. All values are in ppbv.

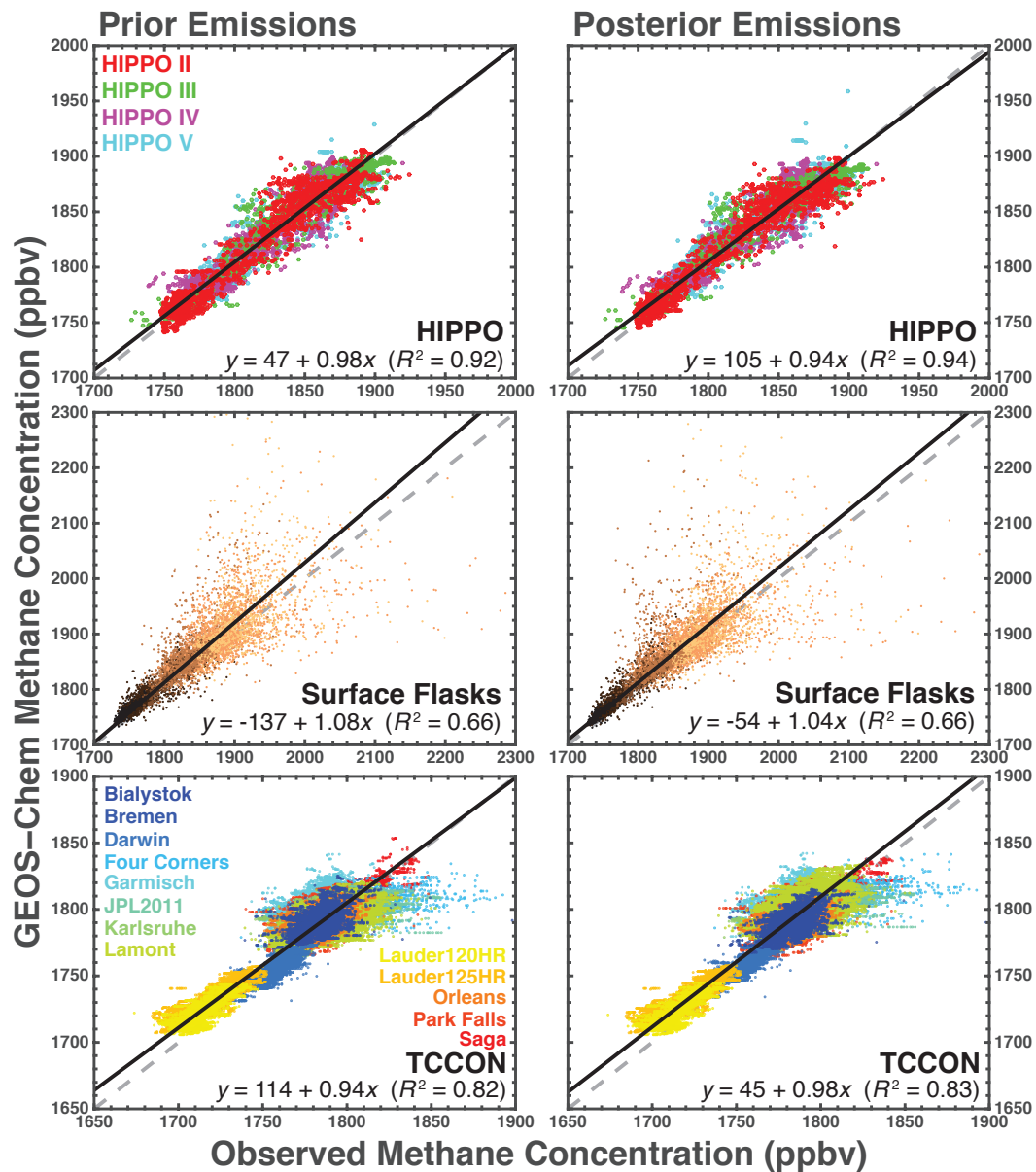


Figure A5. Scatterplot comparison of GEOS-Chem at $4^\circ \times 5^\circ$ resolution to independent observations. Left column uses prior emissions and right column uses posterior emissions. Individual points show comparisons for individual observations, averaged over the GEOS-Chem grid resolution in the case of the aircraft data. The 1 : 1 (dashed) and reduced-major-axis (RMA, solid) regression lines are shown. Summary statistics are also given in Table 2. Different colors correspond to different sites (TCCON), latitudinal bands (flasks), and deployments (HIPPO) shown in Fig. A3.