1 On the progress of the 2015–2016 El Niño event

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10 Abstract

It has been recently reported that the current 2015–2016 El Niño could become "one of the strongest on record". To further explore this claim, we performed the new analysis described in detail in Varotsos et al. (2015) that allows the detection of precursory signals of the strong El Niño events by using a recently developed non-linear dynamics tool. In this context, the analysis of the Southern Oscillation Index time series for the period 1876–2015 shows that the running 2015–2016 El Niño would be rather a "moderate to strong" or even a "strong" event and not "one of the strongest on record", as that of 1997–1998.

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19 **1** Introduction

20 El Niño/La Niña Southern Oscillation (ENSO) is an oceanic-atmospheric quasi-periodic 21 phenomenon with several impacts on climate and weather not only in the tropical Pacific, but 22 in many regions all over the world (Varotsos and Deligiorgi, 1991; Kondratyev and Varotsos, 23 1995a,b; Klein et al., 1999; Xue et al., 2000; Eccles and Tziperman, 2004; Cracknell and 24 Varotsos, 2007, 2011; Lin, 2007; Chattopadhyay and Chattopadhyay, 2011; Efstathiou et al., 25 1998, 2011; Varotsos, 2013; Varotsos et al., 2009a, 2012, 2014a,b). The disastrous effects of 26 the strong ENSO events necessitate their reliable short and long-term prediction (Latif et al., 27 1998; Stenseth et al., 2003; Monks et al., 2009; Hsiang et al., 2011; Cheng et al., 2011; 28 Barnston et al., 2012; Krapivin and Shutko, 2012; Tippett et al., 2012). In this context, 29 Varotsos et al. (2015) presented a new method (see also Varotsos and Tzanis, 2012) for the

detection of precursory signals of the strong El Niño events by using the entropy change in "natural time" (a new time domain, see Varotsos et al., 2002) under time reversal. The analysis of the Southern Oscillation Index (SOI) time series by using this modern method provided significant precursory signals of two of the strongest El Niño events (1982–1983 and 1997–1998).

Very recently, Klein (2015) reported that the running 2015–2016 El Niño could become "one 6 7 of the strongest on record". Furthermore, the Australian Government Bureau of Meteorology 8 (BOM) in their report 9 (http://www.bom.gov.au/climate/enso/archive/ensowrap 20150901.pdf) of 1 September 2015 10 stated that "The 2015 El Niño is now the strongest El Niño since 1997-98" and moreover on 11 29 September 2015 they reported that most international climate models indicate current El 12 Niño (http://www.bom.gov.au/climate/enso/archive/ensowrap 20150929.pdf) "is likely to 13 peak towards the end of 2015" as also reported on 8 October 2015 by the Climate Prediction 14 Center, National Centers for Environmental Prediction, National Oceanic and Atmospheric 15 Administration (NOAA)/National Weather Service 16 (http://www.cpc.ncep.noaa.gov/products/analysis_monitoring/enso_disc_oct2015/ensodisc.pdf).

In this study, we further explore these claims, by applying to the SOI time series the recently proposed analysis by Varotsos et al. (2015). The ability of accurate predictions of such severe natural events, like El Niño, is of crucial importance especially nowadays, where the global annual average temperature in 2015 reached the warmest on record values, which might be associated with the 2015 El Niño event (WMO, 2016).

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23 2 Results and discussion

As mentioned in the previous section, we analyse the SOI time series (Troup, 1965; Power 24 25 and Kociuba, 2011) for the period January 1876 – October 2015 by employing the method 26 described in detail in Varotsos et al. (2015). More specifically, we conduct the analysis of the 27 SOI monthly values by using the dataset, entitled "Monthly SOIPhase 1887 – 1989 Base", 28 (https://www.longpaddock.qld.gov.au/seasonalclimateoutlook/southernoscillationindex/soidatafiles/index.php) 29 derived from the Long Paddock site. It should be clarified that we use the monthly values of 30 SOI, instead of the daily ones, as the latter introduce significant noise due to daily weather 31 patterns variability. It should be noted here that El Niño and La Niña episodes are associated 32 with negative and positive values of the SOI, respectively, and SOI =

1 $10 \times [PA(Tahiti) - PA(Darwin)]/SDD$, where the Pressure Anomaly (*PA*) is the monthly mean 2 minus long-term mean (1887–1989 base period) and *SDD* is the standard deviation of the 3 difference (1887–1989 base period) of mean sea level pressure between Tahiti and Darwin.

4 The method suggested by Varotsos et al. (2015) is based on the entropy change in natural time 5 under time reversal ΔS_i (e.g., see Varotsos et al., 2005, 2007, 2009b; Sarlis et al., 2010, 2011) calculated for a window size of *i* events (SOI monthly values). To this end, Varotsos et al. 6 (2015) converted the original SOI time series to a new one $Q_k = (SOI_k + |min(SOI)|)$, where 7 8 min(SOI) is the minimum value of SOI during the whole study period, keeping the temporal 9 sequence of the events and not considering their time of occurrence. Hence, for each Q_k value we calculate the ratio (χ_k) of the order of its occurrence (k) and the total number (i) of events 10 within the window, i.e. $\chi_k = k/i$. The latter quantity, which replaces the conventional time 11 (t), is natural time χ_k characterizing the k-th event (Varotsos et al., 2002). This way, 12 Varotsos et al. (2015) introduced a new series the members of which are the pairs (χ_k, Q_k) 13 where $Q_k > 0$. Thus, one can define the quantity $p_k = Q_k / \sum_{n=1}^i Q_n$ which can be considered 14 as a probability, since it is positive and satisfies the condition $\sum_{n=1}^{i} p_n = 1$ (Varotsos et al., 15 2011). Under these assumptions, the average values of quantities, which are functions of 16 natural time χ , can be evaluated by $\langle f(\chi) \rangle = \sum_{n=1}^{i} f(\chi_n) p_n$ and the entropy in natural time 17 can be defined by $S = \langle \chi \ln \chi \rangle - \langle \chi \rangle \ln \langle \chi \rangle$ (Varotsos et al., 2005, 2011). The latter quantity 18 19 changes to a value S_{-} if, instead of the true sequence of events, one uses the time-reversed process that is described by $p'_{k} = \hat{T}p_{k} = p_{i-k+1}$, where \hat{T} denotes the time reversal operator in 20 the window of *i* events. The quantity ΔS_i (= S - S_) reveals the breaking of time-symmetry by 21 22 capturing the difference in the dynamics as the system evolves from present to future and 23 vice-versa. In short, it has been shown (e.g., see Varotsos et al., 2007, 2011) that positive 24 values of ΔS_i correspond to a decreasing time-series in natural time, and hence when ΔS_i 25 exceeds a certain threshold this reveals that SOI is approaching at small values indicating El Niño (Varotsos et al., 2015). Varotsos et al. (2015) have also shown (see their Fig. 4) that the 26 27 most useful window size for this purpose is i = 20 events (months). In their prediction scheme, 28 the monthly SOI values for the past 20 months are used for the calculation of ΔS_{20} (see the red 29 crosses in Figs. 1 and 3) and compared with a threshold ΔS_{thres} , which can be determined on

1 the basis of Receiver Operating Characteristics (ROC, see Fawcett, 2006). ROC is a method 2 for the visualization, evaluation, and selection of prediction schemes based on their performance, which is quantified by a plot of the hit rate vs. the false alarm rate obtained by 3 the following procedure applied to the present case. When $\Delta S_{20} \ge \Delta S_{\text{thres}}$, one issues an alarm 4 5 that the value of SOI for the next month will be smaller than or equal to T (see the black 6 broken line in Fig. 2). If this turns out to be true, then we have a true positive prediction. If $\Delta S_{20} < \Delta S_{\text{thres}}$ and the next month's SOI is larger than T, then we have a true negative 7 prediction. All other combinations lead to errors (which are inevitable in stochastic 8 9 prediction), which can be either false positive or false negative predictions. Figure 2 depicts 10 the ROC curve obtained, when using ΔS_{20} as a predictor for the SOI value of the next month 11 with T = -14 (which is the upper limit of the yellow area in Figs. 1 and 3 discussed below). 12 This is a diagram of the hit rate (or True Positive rate, i.e., the number of true positive 13 predictions over all cases with SOI $\leq T = -14$) vs. the false alarm rate (or False Positive rate, 14 i.e., the number of false positive predictions over all cases with SOI > -14) as we vary ΔS_{thres} . 15 A method to estimate an appropriate value of ΔS_{thres} is that of iso-performance lines suggested 16 by Provost and Fawcett (1998, 2001). In this scheme, a line of constant slope m (see the blue 17 line in Fig. 2) is selected on the basis of the relative cost of false positive predictions over the 18 cost of false negative predictions multiplied by the relative frequency of negatives over 19 positives, i.e., see Eq. (1) of Fawcett (2006). As a typical selection we chose m = 1. We fitted ROC points with the red curve (having a simple analytical form $a + b\sqrt{x} + cx^{d}$) and 20 21 determined the point at which the slope was unity. This leads to the ROC point indicated by an arrow in Fig. 2 and corresponds to $\Delta S_{\text{thres}} = 0.0035$ (i.e., a value very close to that 0.00326 22 23 presented in Table 1 of Varotsos et al (2015) for T = -15). Thus, in Figs. 1 and 3 when $\Delta S_{20} \ge$ 24 0.0035 the alarm is set on for the SOI value of the next month.

25 The time progress of the SOI monthly values as well as the entropy change in natural time 26 under time reversal (for the window length i = 20 months) ΔS_{20} are depicted in Fig. 1 (as well 27 as in Fig. 3). Beyond the information gained from the exploration of the ΔS_{20} dynamics and in order to further identify if 2015-2016 El Niño could be characterized as a "very strong" one 28 or even more as "one of the strongest on record", we followed the classification and 29 of 30 characterization the past El Niño events given by BOM 31 (http://www.bom.gov.au/climate/enso/enlist/). The coloured areas in Figs. 1 and 3 represent 32 the mean minimum negative values of SOI along with the 1σ standard deviation bands for the

two cases of "weak, weak to moderate, moderate, moderate to strong" (green band) and
 "strong, very strong" (yellow band) El Niño events.

As can be clearly seen in Fig. 3, the SOI values during the last three months remain in the 3 green band and in the limits of the yellow one, indicating that 2015 El Niño should be rather 4 characterized as a "moderate to strong" or even "strong" event and not "one of the strongest 5 on record", as also shown by comparing with the El Niño events of 1982-1983 and 1997-6 7 1998. Furthermore, the variation of ΔS_{20} during the 2015 El Niño in comparison with 1982– 8 1983 and 1997–1998 El Niño events is not as sharp, confirming that the undergoing El Niño 9 event is not "one of the strongest on record". In order to estimate the extent of this variation, 10 we plot with the black curve in Fig. 4 the probability density function (PDF) of ΔS_{20} obtained from the estimator $f_N(\Delta S_{20}) = \frac{1}{N} \sum_{i=1}^N \frac{1}{b_N} K\left(\frac{\Delta S_{20} - O_i}{b_N}\right)$, where O_i are the observed values of 11 12 ΔS_{20} since the beginning of our study, N is the total number of these observations, the kernel K(x) is non-zero only when |x| < 1 having the value $K(x) = \frac{3}{4}(1-x^2)$ and b_N is related with the 13 standard deviation σ of the observed ΔS_{20} values by $b_N = 10.25\sigma/N^{0.34}$ as suggested by 14 15 Mercik et al. (1999). We observe in Fig. 4 that only rarely ΔS_{20} exceeds the value of 0.02, 16 which can be also verified by the red histogram obtained for ΔS_{20} using the TISEAN package (Hegger et al., 1999) (also plotted in Fig. 4). In the latter histogram, the minimum non-zero 17 18 height is observed in the bar that includes the value $\Delta S_{20} = 0.02$ covering the range up to 19 approximately 0.0205. To detect when ΔS_{20} exceeds the latter value, we plot with blue crosses 20 the time series of ΔS_{20} vs. time, which can be read in the right axis of Fig. 4. We see (blue arrows in Fig. 4) that $\Delta S_{20} > 0.0205$ is observed only in the three strong El Niño events of 21 22 1905-1906, 1982-1983 and 1997-1998. This inequality, however, is not fulfilled in the current 23 case (2015–2016 El Niño), since the currently observed values are close to 0.01, i.e., 24 markedly smaller than the value of 0.0205.

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26 3 Conclusions

Recent reports indicate that 2015–2016 El Niño event could become "one of the strongest on
record" or could be already characterized as "the strongest El Niño since 1997–98". In order
to investigate these assertions, we analyzed the SOI time series for the period January 1876 –

October 2015 by using the method described in Varotsos et al. (2015) based on the entropy
 change in natural time under time reversal. The results obtained indicate that the undergoing
 2015–2016 El Niño event should be rather characterized as a "moderate to strong" or even
 "strong" event and not "one of the strongest on record".

5 **References**

Barnston, A. G., Tippett, M. K., L'Heureux, M. L., Li, S. H., and DeWitt, D. G.: Skill of realtime seasonal ENSO model predictions during 2002–11: is our capability increasing? B.
Am. Meteorol. Soc., 93, 631–651, 2012.

- 9 Chattopadhyay, S. and Chattopadhyay, G.: The possible association between summer
 10 monsoon rainfall in India and sunspot numbers, Int. J. Remote Sens., 32, 891-907, 2011.
- Cheng, Y. J., Tang, Y. M., and Chen, D. K.: Relationship between predictability and forecast
 skill of ENSO on various time scales, J. Geophys. Res., 116, C12006,
 doi:10.1029/2011JC007249, 2011.
- Cracknell, A. P. and Varotsos, C. A.: The Antarctic 2006 ozone hole, Int. J. Remote Sens., 28,
 15 1–2, 2007.
- Cracknell, A. P. and Varotsos, C. A.: New aspects of global climate-dynamics research and
 remote sensing, Int. J. Remote Sens., 32, 579–600, 2011.
- Eccles, F. and Tziperman, E.: Nonlinear effects on ENSO's period, J. Atmos. Sci., 61, 474–
 482, 2004.
- Efstathiou, M., Varotsos, C., and Kondratyev, K. Y.: An estimation of the surface solar
 ultraviolet irradiance during an extreme total ozone minimum, Meteorol. Atmos. Phys., 68,
- 22 171–176, 1998.
- Efstathiou, M. N., Tzanis, C., Cracknell, A. P., and Varotsos, C. A.: New features of land and
 sea surface temperature anomalies, Int. J. Remote Sens., 32, 3231–3238, 2011.
- 25 Fawcett, T.: An introduction to ROC analysis, Pattern Recogn. Lett., 27, 861–874, 2006.
- 26 Hegger, R., Kantz, H. and Schreiber, T.: Practical implementation of nonlinear time series
- 27 methods: The TISEAN package, Chaos, 9, 413-435, 1999.
- Hsiang, S. M., Meng, K. C., and Cane, M. A.: Civil conflicts are associated with the global
 climate, Nature, 476, 438–441, 2011.

| 1 | Klein, K.: NOAA predicts strong El Niño, Eos, 96, doi:10.1029/2015EO035535, 2015. |
|--|--|
| 2 3 | Klein, S. A., Soden, B. J., and Lau, N. C.: Remote sea surface temperature variations during ENSO: Evidence for a tropical atmospheric bridge, J. Climate, 12, 917–932, 1999. |
| 4 5 | Kondratyev, K. Y. and Varotsos, C.: Atmospheric greenhouse effect in the context of global climate change, Il Nuovo Cimento C, 18, 123–151, 1995a. |
| 6 7 | Kondratyev, K. Y. and Varotsos, C. A.: Volcanic-eruptions and global ozone dynamics, Int. J. Remote Sens., 16, 1887–1895, 1995b. |
| 8 9 | Krapivin, V. F. and Shutko, A. M.: Information technologies for remote monitoring of the environment, Springer/Praxis, Chichester, UK, 2012. |
| 10 11 12 | Latif, M., Anderson, D., Barnett, T., Cane, M., Kleeman, R., Leetmaa, A., O'Brien, J., Rosati, A., and Schneider, E.: A review of the predictability and prediction of ENSO, J. Geophys. Res., 103, 14375–14393, 1998. |
| 13 14 | Lin, JL.: Interdecadal variability of ENSO in 21 IPCC AR4 coupled GCMs, Geophys. Res. Lett., 34, L12702, doi:10.1029/2006GL028937, 2007. |
| 15 16 | Mercik, S., Weron, K., and Siwy, Z.: Statistical analysis of ionic current fluctuations in membrane channels, Phys. Rev. E, 60, 7343–7348, 1999. |
| 17 18 19 20 21 22 23 24 25 26 27 28 | Monks, P. S., Granier, C., Fuzzi, S., Stohl, A.,Williams, M. L., Akimoto, H., Amann, M., Baklanov, A., Baltensperger, U., Bey, I., Blake, N., Blake, R. S., Carslaw, K., Cooper, O. R., Dentener, F., Fowler, D., Fragkou, E., Frost, G. J., Generoso, S., Ginoux, P., Grewe, V., Guenther, A., Hansson, H. C., Henne, S., Hjorth, J., Hofzumahaus, A., Huntrieser, H., Isaksen, I. S. A., Jenkin, M. E., Kaiser, J., Kanakidou, M., Klimont, Z., Kulmala, M., Laj, P., Lawrence, M. G., Lee, J. D., Liousse, C., Maione, M., McFiggans, G., Metzger, A., Mieville, A., Moussiopoulos, N., Orlando, J. J., O'Dowd, C. D., Palmer, P. I., Parrish, D. D., Petzold, A., Platt, U., Pöschl, U., Prévôt, A. S. H., Reeves, C. E., Reimann, S., Rudich, Y., Sellegri, K., Steinbrecher, R., Simpson, D., ten Brink, H., Theloke, J., van der Werf, G. R., Vautard, R., Vestreng, V., Vlachokostas, Ch., and von Glasow, R.: Atmospheric composition change – global and regional air quality, Atmos. Environ., 43, 5268–5350, 2009. |
| 29 30 | Power, S. B. and Kociuba, G.: The impact of global warming on the Southern Oscillation Index, Clim. Dynam., 37, 1745–1754, 2011. |
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| 1 | Provost, F. and Fawcett, T.: Robust classification systems for imprecise environments, i | in: |
|---|--|-----|
| 2 | Proceedings of the AAAI-98, Menlo Park, CA, 706–713, 1998. | |

- Provost, F. and Fawcett, T.: Robust classification for imprecise environments, Mach. Learn.,
 42, 203–231, 2001.
- Sarlis, N. V., Skordas, E. S., and Varotsos, P. A.: Nonextensivity and natural time: The case
 of seismicity, Phys. Rev. E, 82, 021110, doi:10.1103/PhysRevE.82.021110, 2010.
- 7 Sarlis, N. V., Skordas, E. S., and Varotsos, P. A.: The change of the entropy in natural time
- 8 under time-reversal in the Olami–Feder–Christensen earthquake model, Tectonophysics,
 9 513, 49–53, 2011.
- 10 Stenseth, N. C., Ottersen, G., Hurrell, J. W., Mysterud, A., Lima, M., Chan, K. S., Yoccoz, N.

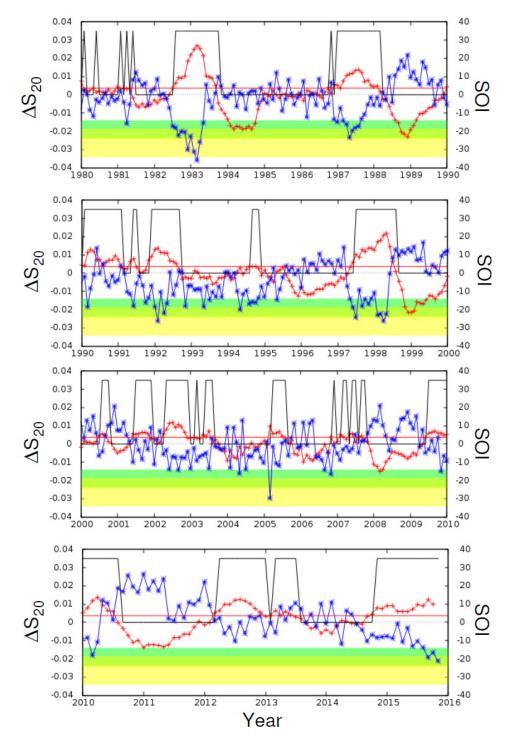
11 G., and Adlandsvik, B.: Studying climate effects on ecology through the use of climate

12 indices: the North Atlantic Oscillation, El Niño Southern Oscillation and beyond, P. Roy.

- 13 Soc. Lond. B Bio., 270, 2087–2096, 2003.
- Tippett, M. K., Barnston, A. G., and Li, S. H.: Performance of recent multimodel ENSO
 forecasts, J. Appl. Meteorol. Clim., 51, 637–654, 2012.
- 16 Troup, A. J.: The Southern Oscillation, Q. J. Roy. Meteor. Soc., 91, 490–506, 1965.
- Varotsos, C. A.: The global signature of the ENSO and SST-like fields, Theor. Appl.
 Climatol., 113, 197–204, 2013.
- 19 Varotsos, C. A. and Deligiorgi, D. G.: Sea-surface temperature and southern oscillation signal
- 20 in the upper stratosphere-lower mesosphere, Int. J. Climatol., 11, 77–83, 1991.
- Varotsos, C. A. and Tzanis, C.: A new tool for the study of the ozone hole dynamics over
 Antarctica, Atmos. Environ., 47, 428–434, 2012.
- Varotsos, C., Efstathiou, M., and Tzanis, C.: Scaling behaviour of the global tropopause,
 Atmos. Chem. Phys., 9, 677-683, 2009a.
- 25 Varotsos, C. A., Cracknell, A. P., and Tzanis, C.: The exceptional ozone depletion over the
- Arctic in January-March 2011, Remote Sens. Lett., 3, 343–352, 2012.
- Varotsos, C., Christodoulakis, J., Tzanis, C., and Cracknell, A. P.: Signature of tropospheric
 ozone and nitrogen dioxide from space: A case study for Athens, Greece, Atmos. Environ.,
 89, 721-730, 2014a.
- *29 89, 721-730, 2014a.*

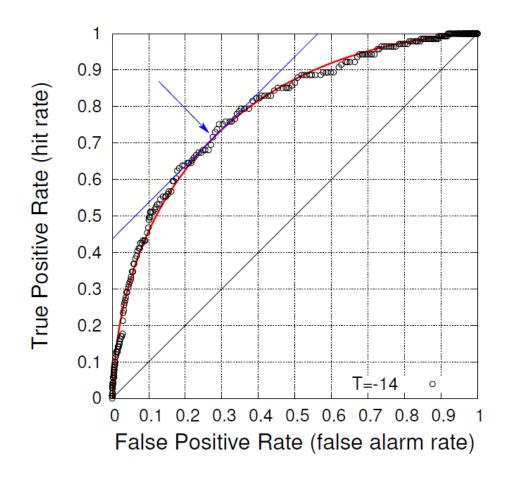
- Varotsos, C. A., Franzke, C. L. E., Efstathiou, M. N., and Degermendzhi, A. G.: Evidence for
 two abrupt warming events of SST in the last century, Theor. Appl. Climatol., 116, 51–60,
 2014b.
- 4 Varotsos, C. A., Tzanis, C., and Cracknell, A. P.: Precursory signals of the major El Niño
 5 Southern Oscillation events, Theor. Appl. Climatol., doi:10.1007/s00704-015-1464-4,
 6 online first, 2015.
- Varotsos, P. A., Sarlis, N. V., and Skordas, E. S.: Long-range correlations in the electric
 signals that precede rupture, Phys. Rev. E, 66, 011902, doi:10.1103/PhysRevE.66.011902,
 2002.
- Varotsos, P. A., Sarlis, N. V., and Skordas, E. S.: Detrended fluctuation analysis of the
 magnetic and electric field variations that precede rupture, Chaos, 19, 023114,
 doi:10.1063/1.3130931, 2009b.
- Varotsos, P. A., Sarlis, N. V., Tanaka, H. K., and Skordas, E. S.: Some properties of the
 entropy in the natural time, Phys. Rev. E, 71, 032102, doi:10.1103/PhysRevE.71.032102,
 2005.
- Varotsos, P. A., Sarlis, N. V., Skordas, E. S., and Lazaridou, M. S.: Identifying sudden cardiac
 death risk and specifying its occurrence time by analyzing electrocardiograms in natural
 time, Appl. Phys. Lett., 91, 064106, doi:10.1063/1.2768928, 2007.
- 19 Varotsos, P. A., Sarlis, N. V., and Skordas, E. S.: Natural Time Analysis: The new view of
- time. Precursory Seismic Electric Signals, Earthquakes and other Complex Time-Series,
 Springer-Verlag, Berlin Heidelberg, 2011.
- WMO (World Meteorological Organization), Press Release No. 2, 2016, available at:
 https://www.wmo.int/media/content/2015-hottest-year-record.
- Xue, Y., Llewellyn-Jones, D.T., Lawrence, S. P., and Mutlow, C. T.: On the Earth's surface
 energy exchange determination from ERS satellite ATSR data: Part 3. Turbulent heat flux
 on open sea, Int. J. Remote Sens., 21, 3427–3444, 2000.

| 1 | Figure captions |
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| 2 | |
| 3 | Figure 1. The entropy change ΔS_{20} in natural time for the window length $i = 20$ months (red |
| 4 | line, left scale) along with SOI monthly values (blue line, right scale) for the period January |
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| 14 | Figure 3. As in Fig. 1, but only for the 1982–1983, 1997–1998 (the two strongest in the last |
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| 19 | crosses, right scale) along the vertical axis. The arrows indicate when ΔS_{20} exceeds 0.0205 |
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Figure 1. The entropy change ΔS_{20} in natural time for the window length i = 20 months (red line, left scale) along with SOI monthly values (blue line, right scale) for the period January 1980 – October 2015. The alarm is set on (black line), when ΔS_{20} exceeds the threshold value $\Delta S_{\text{thres}} = 0.0035$.



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Figure 2. The hit rate vs. false alarm rate when using ΔS_{20} as a predictor for the SOI value of the next month. The ROC point indicated by the arrow has been selected so that the slope of the tangent of the analytical fitting of the ROC points indicated by the red curve has unit slope and hence it corresponds to the m = 1 iso-performance line of the ROC space (e.g., see Fawcett, 2006; Provost and Fawcett, 1998, 2001).

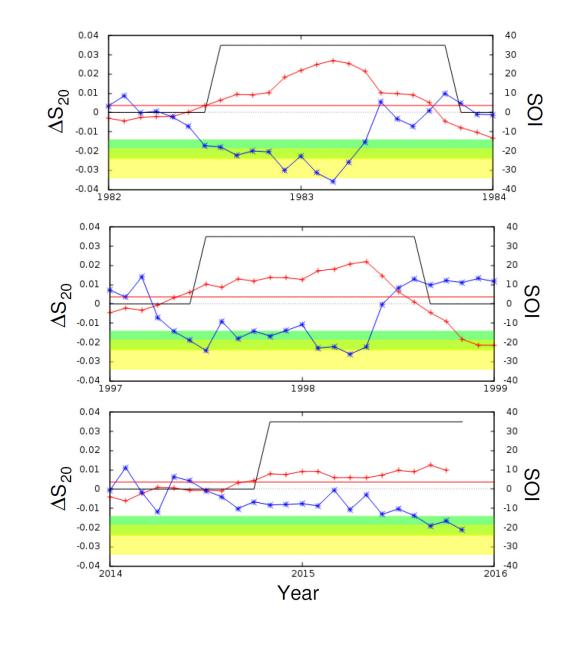


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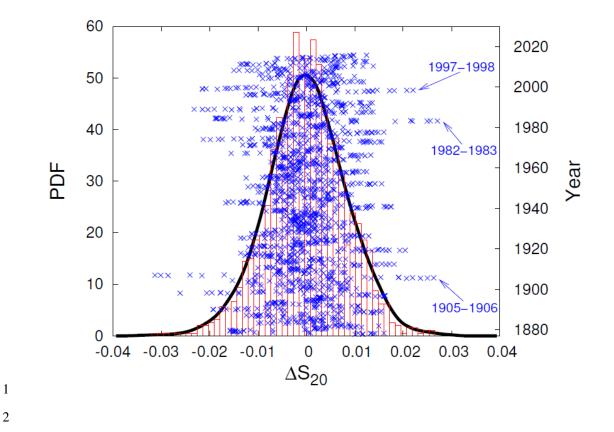


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